

**Planetary analog sites in the Indian subcontinent and the Indian Ocean:
Underexplored environments suited for astrobiological and space research**

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20 Abstract

21 The central objective of astrobiology is to evaluate the habitability of other planetary bodies and
22 given the high cost and logistical complexity of space missions, preliminary investigations of
23 planetary analog sites are a critical step in supporting and de-risking future exploration efforts.
24 Planetary analog sites are extreme environments on Earth that exhibit one or more environmental,
25 geological, geochemical, or biological characteristics that are analogous to conditions expected on
26 other planetary bodies. Understanding how life persists in planetary analog sites is essential for
27 advancing our knowledge of extraterrestrial habitability. They provide valuable opportunities to
28 study life's resilience and to test life-detection instruments in realistic settings. In this review, we
29 present the first comprehensive synthesis of over 30 planetary analog field sites located across the
30 Indian subcontinent and the surrounding Indian Ocean region. These include high-altitude glaciers,
31 alkaline lakes, hypersaline basins, hot springs, and cold desert ecosystems. Although these
32 environments remain largely underexplored in the context of astrobiology, they exhibit strong
33 environmental parallels to planetary conditions and offer significant potential for advancing
34 microbial ecology, biosignature detection, and geobiological research. As the scientific community
35 prepares for the next generation of deep space missions focused on the search for life beyond Earth,
36 we call for a broader geographical and conceptual inclusion of analog sites. By highlighting the
37 environmental diversity and scientific value of these under-characterized regions in South Asia and
38 their marine periphery, this review contributes a vital and previously underrepresented perspective to
39 planetary analog research. This review fills a critical gap in the global planetary analog landscape and
40 aims to stimulate interdisciplinary investigations across planetary science, microbiology, and life-
41 detection technology development.

42 Introduction

43 Astrobiology explores fundamental questions about the origin, evolution, distribution, and future of
44 life on Earth and beyond ([Chyba and Hand 2005](#)). The search for life outside Earth begins with
45 identifying and characterizing potentially habitable environments within our Solar System bodies
46 such as Mars, Europa, and Enceladus, as well as on exoplanets orbiting distant stars ([McKay 2022](#);
47 [Kane et al. 2021](#)). This search is informed by our understanding of how life first emerged and
48 evolved on Earth. Early life thrived in environments we now call “extreme,” is a designation that
49 reflects an anthropocentric perspective, since these conditions represent physiologically normal
50 habitats for the microorganisms that inhabit them ([Rothschild and Mancinelli 2001](#); [Horneck et al.](#)
51 [2010](#); [Merino et al. 2019](#)). In extreme environments, microbes can tap into inorganic compounds
52 such as iron, sulfur, or hydrogen for energy, in absence of sunlight and underscoring the possibility
53 of life in dark subsurface habitats beyond Earth ([Canfield et al. 2006](#)). Moreover, microbes can
54 exploit a wide range of electron donors and acceptors, enabling them to drive diverse metabolic
55 pathways and thrive across varied environmental conditions ([Froelich et al. 1979](#)). Physicochemical
56 extremes, including high radiation, hypersalinity, pressure variations, and nutrient-poor terrains, still
57 persist on Earth today, delineating the outer limits of habitability ([Merino et al. 2019](#); [Coleine and](#)
58 [Delgado-Baquerizo 2022](#)). Such environments serve as planetary analogues, offering natural
59 laboratories to investigate microbial survival, biosignature preservation, and the potential for life on

other worlds. They also play a critical role in developing and validating life-detection instruments, refining scientific methodologies, and supporting mission planning through terrain and environmental simulations ([Léveillé 2009](#); [Martins et al. 2017](#)). In recent years, the concept of planetary analogues has expanded beyond purely “geological analogues” to include “functional analogues.” These may involve natural sites or engineered platforms selected for their relevance to specific scientific or operational goals, such as petrology, mineralogy, biological processes, or engineering compatibility ([Foucher et al. 2021](#)). This broader classification enables a more strategic integration of analogue environments across mission phases; from instrumentation development to in-situ exploration and post-mission data interpretation

International space agencies have prioritized a range of terrestrial field sites as astrobiological analogues (Supplementary Table 1) for planetary bodies such as Mars, Venus, Lunar, and Icy Moons (Europa and Enceladus). For Mars, Venus, and other rocky planets, hyperarid deserts and volcanic terrains provide valuable parallels. The Atacama Desert in Chile, one of the driest places on Earth, mirrors the UV-intense, desiccated surface of Mars and is critical for studying microbial survival and biosignature stability ([Azua-Bustos et al. 2022](#); [Lebre et al. 2017](#); [Visscher et al. 2020](#)). Likewise, Death Valley ([Farr 2004](#); [Douglas et al. 2008](#)), the Tibetan Plateau ([Liu et al., 2022](#); [Sun et al., 2018](#)), and the McMurdo Dry Valleys in Antarctica offer Mars-like extremes of aridity, cold, and low atmospheric pressure ([Salvatore and Levy 2021](#)). Hydrothermal and volcanic sites, such as the Iberian Pyritic Belt in Spain ([Amils et al. 2007](#); [2014](#); [Gómez et al. 2011](#)), the Dallol geothermal field in Ethiopia ([Cavalazzi et al. 2019](#)), Icelandic volcanoes ([Voigt et al., 2025](#); [Fagents and Thordarson, 2007](#)), and the basalt flows of Kīlauea, Hawaii ([Bergsten et al., 2021](#); [D’Incecco et al., 2024](#); [Seelos et al., 2010](#)), recreate acidic and mineral-rich terrains comparable to past Martian hydrothermal systems and Venusian volcanism. The Manicouagan impact crater in Canada further contributes insights into impact-generated habitats relevant to Mars and icy bodies ([Spray et al. 2010](#); [Cloutis et al. 2015](#)). Svalbard in Norway provides permafrost landscapes and logistical isolation that make it an effective analog for Mars surface exploration and lunar habitation scenarios ([Kołodziejczyk et al., 2025](#); [Hauber et al., 2011](#)). For icy moons, subglacial and deep-sea hydrothermal environments offer compelling parallels. Lake Vostok in Antarctica ([Cassaro et al. 2021](#)), the Strýtan hydrothermal vents in Iceland ([Twing et al. 2022](#)), and seafloor systems such as Gakkel Ridge ([Ramirez-Llodra et al. 2023](#)), Lost City ([Amador et al. 2017](#)), Juan de Fuca ([Kelley et al. 2012](#); [Osinski et al. 2006](#)), and the Mariana Trench ([Preston and Dartnell 2014](#)) simulate energy-rich, water–rock interactions in subsurface oceans where life may persist ([Gaidos et al. 1999](#); [Barge and White 2017](#); [Martin and McMinn 2018](#); [German et al. 2022](#)). Despite global progress in analogue research, geologically diverse regions remain underexplored. The Indian subcontinent and adjoining Indian Ocean basin, shaped by complex tectonics and unique environmental extremes ([Valdiya 2016](#)), represent promising yet overlooked additions to the planetary analogue inventory. Incorporating these sites would broaden the global scope of astrobiology, expand analogue functions, and deepen our understanding of habitability across a wider spectrum of planetary environments.

The Indian subcontinent and the adjoining Indian Ocean, with their tectonic complexity, encompass diverse geological settings that offer valuable opportunities for testing hypotheses on the origin and resilience of life in extraterrestrial environments (Figure 1 and 2). Yet, despite hosting numerous

extreme habitats, these regions remain underexplored in astrobiological research. This gap stems from limited integration of regional geology with planetary science, fragmented interdisciplinary collaborations, and significant practical challenges. Many sites are remote and difficult to access, while sustained field-based studies are constrained by limited economic resources. Additionally, anthropogenic activity and contamination in some areas complicate the preservation of pristine conditions for analogue research. However, these geological field sites retain high relevance, as their tectonic, mineralogical, and geochemical attributes preserve natural extremes that provide valuable parallels to planetary environments (Supplementary Table 2,3, and 4; Figure 1 and 2). The tectonic collision between the Indian and Eurasian Plates, around 50 Ma ([Molnar and Tapponnier 1977](#); [Jain 2014](#); [Wu et al. 2023](#)), gave rise to the Himalayas and exposing stratigraphic sequences spanning from the Archean geological time (4 Ga), including some of the oldest continental crusts preserved in the Dharwar and Singhbhum cratons ([Dey and Moyen 2020b](#)) and the ophiolitic sequences ([Bhat et al. 2022](#)), ideal for Early Earth investigations. High-altitude glacial systems, such as Ladakh, Sikkim, and the Hindu Kush-Himalayan ranges, mimics the Martian and/or early Earth Environments ([Pandey et al. 2020](#); [Schopf et al. 2008](#); [Armstrong 2010](#)). Dinodhar hill (Kachchh) and Alaldari Valley (Maharashtra) offer parallels to fluvial activity on Martian surfaces ([Chavan et al. 2022](#)). Hyper-arid zones, such as the Thar Desert may simulate Martian surface desiccation and dust dynamics ([Khormali and Monger 2020](#); [Sarkar et al. 2019](#)). The hypersaline Sambhar lake ([Sarkar et al. 2023](#); [Sinha and Raymahashay 2004a](#)) and the Lonar soda lake ([Paul Antony et al. 2013](#)) present analogs to evaporitic basins on Mars ([Mancinelli et al. 2004](#)). Additionally, geothermal springs scattered across the Deccan Traps and the Himalayan region ([Das et al. 2022](#); [Tamburello et al. 2022](#)), along with the deep-sea hydrothermal vents spread across the Indian Ocean ([Dick 2019a](#)), may provide windows into subsurface microbial life processes relevant to ocean worlds.

This review aims to foreground these underexplored Indian field sites within global analogue research by systematically evaluating their planetary and functional relevance. We highlight their capacity to inform biosignature detection, expand the range of habitability scenarios, and support the development of mission-relevant technologies. Establishing the Indian subcontinent and adjoining Indian Ocean as integral components of the planetary analogue framework will broaden astrobiology's geographic scope and enrich its experimental and operational dimensions.

The Indian Plate

The Indian Plate, composed of Archean cratons, Proterozoic mobile belts, Phanerozoic sedimentary basins, large igneous provinces, and tectonically active regions such as the Himalayas, preserves a geological record that spans from the Archean to the present ([Chaudhuri et al. 2018](#)), hosting extreme environments as listed in Supplementary Table 2 (Figure 3, 4). The Archean cratons (dating between 3.6 and 2.5 Ga), including the Dharwar, Singhbhum, Bastar, Aravalli, and Bundelkhand cratons, are dominated by rocks that preserve the earliest continental crust and provide critical insights into early Earth processes. Stromatolites and banded iron formations within these cratons ([Beukes et al. 2008](#); [Mukhopadhyay et al. 2025](#); [Dey and Moyen 2020a](#)) may have records of microbial life in an anoxic

world. The younger Proterozoic mobile belts (dated 2.5 - 0.5 Ga), include the Aravalli–Delhi Belt, the Eastern Ghats Mobile Belt, and the Central Indian Tectonic Zone, that document ocean closures, continental collisions, crustal reworking, Great Oxidation Event (2.4 Ga), and the Neoproterozoic Oxygenation Event (0.5 billion years ago) ([Kaiho et al. 2024](#); [Olejarz et al. 2021](#); [Chen et al. 2022](#)). Glaciogenic deposits within these belts preserve evidence of microbial life survival during Snowball Earth episodes ([Hoffman et al. 2017](#)). In addition, Precambrian alluvial successions, which formed before the appearance of land plants, retains microbial textures and mineralogical signatures that illuminate early surface processes and offer potential extraterrestrial analogs (Chakraborty et al. 2022). From an astrobiological perspective, these Precambrian records are particularly valuable because they reflect microbial ecosystems under low-oxygen atmospheres, conditions that may resemble those of other habitable planets or moons.

The Phanerozoic era (beginning about 0.5 Ga to the present) brought extensive sedimentation across Gondwana basins. Beginning in the Late Carboniferous and Jurassic, these basins preserve fluvial sandstones, coal beds, and fossiliferous strata that chronicle terrestrial ecosystem evolution during Gondwana’s breakup ([Chakraborty et al. 2019](#)). Along the northern margin of the plate, the Tethyan Himalayan Basin accumulated marine sediments from the Cambrian to the Eocene, recording the history of the vanished Tethys Ocean ([Liu and Einsele 1994](#); [Jiang et al. 2016](#)). Fossils and organic molecules in these sediments, including kerogen, hopanes, and steranes ([Brocks and Summons 2003](#)), provide benchmarks for biosignature interpretation. Phanerozoic strata also record redox variability, ranging from anoxic to sulfidic to oxic conditions, which directly influences biosignature preservation ([Sperling et al. 2021](#)). Their lacustrine ([Michalski et al. 2022](#)), fluvial ([Cardenas et al. 2022](#)), deltaic ([Di Achille and Hynek 2010](#); [De Toffoli et al. 2021](#)), and evaporitic environments ([Kite and Conway 2024](#)) serve as planetary analogs for sedimentary environments identified on Mars. While Phanerozoic basins refine biosignature detection methods in post-vegetation ecosystems, Precambrian archives remain the closest parallels to microbial worlds that may exist on other planets.

In Cenozoic era, the Deccan Traps (emplaced about 66 Ma) form one of Earth’s largest flood basalt provinces and are linked to mantle plume activity and the end-Cretaceous mass extinction ([Sen et al. 2009](#); [Beane et al. 1986](#); [Mishra et al. 2024](#)). Deccan basalts, with their basaltic stratigraphy, alteration patterns, and mineralogy, serve as analogs for Martian terrains such as Mawrth Vallis ([Craig et al., 2017](#); [Bhattacharya et al., 2016](#); [Greenberger et al., 2012](#)). Their porosity and fracture

170 networks host endolithic and subsurface microbial communities that exploit iron- and sulfur-based
171 chemoautotrophy ([Dutta et al. 2019; 2018; Jungbluth et al. 2016](#)). Such systems offer parallels to
172 potential subsurface life on Mars, where rock–water interactions and redox gradients may sustain
173 habitability.

174 Altogether, the Indian Plate encompasses extreme environments shaped by tectonics, climate, and
175 deep time. From Archean cratons and Proterozoic mobile belts to Phanerozoic basins and Cenozoic
176 flood basalts, these natural laboratories preserve biosignatures, track planetary environmental
177 transitions, and provide Earth analogs for diverse extraterrestrial settings, from early Mars to the
178 subsurface oceans of icy moons such as Europa and Enceladus. Studying biosignatures in these
179 terrestrial settings will allow us to constrain what types of biosignatures might form, survive, or
180 degrade in Martian equivalents, particularly in clay-rich ([Bishop et al. 2018; Gainey et al. 2017;](#)
181 [Ehlmann et al. 2008](#)) or iron-rich sediments ([Kizovski et al. 2025; Ehlmann and Edwards 2014](#)).

182 **Soda and hypersaline Lakes**

183 Soda lakes, characterized by high concentrations of sodium, carbonate, and bicarbonate ions, result in
184 highly alkaline waters. These systems are often endorheic, with no fluvial outflow accumulating salts
185 over geological time with persistent high-pH environments on Earth ([Boros and Kolpakova 2018;](#)
186 [Kempe and Kazmierczak 2011](#)). These lakes supports diverse and robust microbial communities
187 adapted to extreme alkalinity ([Dimitry Y Sorokin et al. 2015; Jones et al. 1998; Sorokin et al. 2014;](#)
188 [Sorokin and Kuenen 2005; Zorz et al. 2019](#)). Globally, they serve as natural laboratories for early
189 Earth and potentially habitable environments on other planets ([Cohen et al. 2024; Ventosa 2004](#)).
190 Some of the most well-known soda lakes include Lake Magadi and Lake Natron in East Africa
191 ([Deocampo and Renaut 2016](#)), Big Soda and Mono Lake in California ([Cloern et al. 1983; Honke et](#)
192 [al. 2019; Kharaka et al. 1984](#)), the Kulunda Steppe lakes in Russia ([Meyer et al. 2008](#)), and lakes of
193 the Cariboo plateau in Canada ([Renaut 1990](#)). They support dense populations of alkaliphilic and
194 halophilic microbes, including cyanobacteria, sulfur-oxidizing bacteria, and alkaliphilic archaea
195 ([Jeilu et al. 2022; Cloern et al. 1983; Humayoun et al. 2003; Rojas et al. 2018; Foti et al. 2008;](#)
196 [Dimitry Y. Sorokin et al. 2015; Zorz et al. 2019](#)). Many of these microbes thrive on anaerobic
197 metabolisms such as fermentation, methanogenesis, denitrification, metal respiration, arsenate
198 reduction. The presence of layered microbial mats, carbonate precipitates, and microbialites in these
199 systems provides insights into potential biosignature preservation, making soda lakes key field sites
200 for astrobiological studies.

In the Indian subcontinent, soda lakes are less extensively studied. Notable examples include Lonar Crater Lake, formed by a meteorite impact ([Fredriksson et al. 1973](#)). It hosts microbial communities adapted to both alkaline and saline conditions, including cyanobacteria, purple sulfur bacteria, and archaea ([Paul Antony et al. 2013](#)). The combination of high pH, basaltic geology, and impact-driven hydrology makes Lonar Lake a key planetary analog for Mars ([Pandey et al. 2019](#)), where alkaline brines and volcanic terrains are thought to occur. Another significant site is Sambhar Lake in Rajasthan, which is India's largest inland saline lake and exhibits moderate alkalinity ([Sinha and Raymahashay 2004b](#)). Although it is primarily known for its salt production and avian biodiversity, the lake's fluctuating hydrology and hypersaline conditions make it a potentially interesting site for studying microbial adaptation in alkaline brines ([Sahay et al. 2012](#); [Pal et al. 2020](#)). Both lakes are valuable for exploring alkaline geochemistry, microbial adaptation, prebiotic chemistry, and biosignature preservation in extreme saline–alkaline settings.

Dust Storms of Arid Desert

The hot, arid regions of the Thar and Cholistan Deserts serve as valuable terrestrial analogues for planetary research, particularly for understanding Martian surface processes including dust storms ([Wang and Richardson 2015](#); [Leovy et al. 1973](#)). These deserts are characterized by intense solar radiation, extreme temperature fluctuations and low water availability. In Thar desert, the frequent occurrence of mineral dust storms, reducing visibility, modifying sediment textures, and driving extensive aeolian transport ([Middleton et al. 1986](#); [Jain et al. 2025](#)). These airborne dust influences planetary climate, habitability, and the remote detection of biosignatures (Boutle et al. 2020). These dust storms provide an opportunity to evaluate how dust loading, particle composition, and scattering properties affect remote sensing signals, thereby refining strategies for biosignature detection on Mars and other planetary surfaces. While the physical dynamics of dust storms are similar on Earth and Mars, their compositions differ. Rajasthan dust contains up to 10% iron oxides ([Mishra and Tripathi, 2008](#)), whereas Martian dust is estimated at about 22% ([Berger et al., 2016](#)), due to higher iron content. Dust storms may also transport microbial cells, as demonstrated in other desert systems ([Yamaguchi et al. 2012](#); [González-Toril et al. 2020](#)), raising questions about microbial survival during atmospheric transit and even the potential for interplanetary dispersal ([Barberán et al., 2015](#); [Amato et al., 2023](#)). However, airborne microbial communities in these deserts remain unstudied, with most research focused on soils, dunes, and rock surfaces ([Rao et al., 2016](#); [Parihar et al., 2022](#); [Fatima et al., 2019](#)). Developing predictive models and deploying automated samplers before dust events, or collecting post-storm deposited dust, may provide a way forward in addressing this gap.

Microbial communities in the Thar and Cholistan exhibit adaptations such as spore formation, UV-protective pigments, DNA repair mechanisms, and metabolic flexibility, which enable survival under nutrient-poor and water-limited conditions. These traits, along with the preservation of organic compounds in dry sediments, make the region a natural laboratory for refining life-detection strategies relevant to Mars and other planetary bodies. Lastly, the geomorphology of the Thar, with its active dune fields, evaporite crusts, and silicate-rich sediments, also provides a natural test bed for rover mobility, autonomous navigation, and remote sensing systems designed for planetary exploration.

Natural caves and unused mines

Natural cave environments provide stable temperatures, radiation shielding, and protection from extreme weather and micrometeorite impacts, making them valuable analogues for Martian habitats ([Blank, 2018](#); [Boston et al., 2001](#); [Léveillé & Datta, 2010](#)). Many terrestrial caves form in karst landscapes, where groundwater dissolves soluble rocks such as limestone, dolomite, and gypsum ([Waele and Gutiérrez 2022](#)). On Mars and the Moon, extensive cave networks, including lava tubes and pit craters, have been identified, with some lunar caves maintaining stable temperatures suitable for habitation ([Sauro et al. 2020](#)). Earth-based caves thus serve as planetary analogues and provide opportunities to investigate biosignature preservation, microbial–mineral interactions, and the ecological impacts of long-term darkness. They host diverse extremophilic microorganisms, many relying on chemolithoautotrophy and minimal nutrients ([Gabriel and Northup, 2013](#); [Zhu et al., 2022](#); [Turrini et al., 2024](#)). Lava tubes, in particular, are also useful for testing robotic exploration technologies ([Morrell et al. 2024](#)). Globally, several caves hold astrobiological significance. The Movile Cave in Romania, an oxygen-poor ecosystem, sustains microbial communities based on sulfur and ammonium chemolithotrophy as well as methanogenesis, methanotrophy, and methylotrophy ([Hutchens et al., 2004](#); [Chen et al., 2009](#); [Wischer et al., 2015](#)). Lechuguilla Cave in New Mexico harbors antibiotic-resistant microbes preserved in pristine conditions, providing insights into the evolution of the deep biosphere ([Cunningham et al. 1995](#)). The Frasassi Caves in Italy support sulfur-cycling microbial mats that exemplify metabolisms relevant to extreme environments ([Macalady et al. 2006; 2007](#)). In the Indian subcontinent, Borra Caves, Krem Phyllut, Mawsmai, and Kotumsar Cave (Chhattisgarh) offer promising settings for astrobiological research (Baskar and Baskar 2022; Baskar et al. 2009; 2011).

Beyond natural caves, unused or decommissioned mines provide valuable subsurface laboratories for astrobiology. Deep underground environments replicate conditions relevant to planetary interiors, with high pressure, limited nutrients, and absence of sunlight. Around the world, several mines have been repurposed: the Sanford Underground Research Facility in South Dakota, converted from the Homestake gold mine, now hosts experiments in physics and subsurface biology ([Heise and J. 2015](#); [Osburn et al. 2014](#); [Jangir et al. 2019](#); [Osburn et al. 2019](#); [Rowe et al. 2021](#)); the Boulby Underground Laboratory in the UK supports planetary science and microbiology research in a potash mine ([Wadsworth et al. 2020](#); [Cockell et al. 2019](#); [Payler et al. 2019](#)); the Canadian Underground Research Laboratory revealed biofilm-associated minerals in granite ([Brown et al. 1994](#)); the Soudan mine has yielded diverse iron-oxidizing and -reducing microbes ([Hsu et al. 2024](#); [Badalamenti et al. 2016](#); [Edwards et al. 2006](#)); and deep mines in South Africa have produced groundbreaking discoveries of microbial taxa thriving in high-pressure conditions ([Gihring et al. 2006](#); [Onstott et al. 1997](#)). Similar opportunities also exist in India. The Kolar Gold Fields in the Dharwar Archean craton ([Siddaiah and Rajamani 1989](#); [Reddy et al. 2017](#)), sections of the Paleoproterozoic regions hosted in Jaduguda uranium mines in the Singhbhum shear zone ([Pal et al. 2011](#)), the Zawar zinc mines in Rajasthan, and the Malanjkhand copper mine in Madhya Pradesh ([Pandey et al. 2007](#); [Equeenuddin et al. 2017](#); [Sikka 1989](#)) expose diverse lithologies and complex geochemical settings. These environments are ideal for investigating extremophile microbial communities, biogeochemical cycling, and biosignature preservation. Repurposing such sites as underground laboratories would enable studies of water–rock–microbe interactions, radiation shielding strategies, and life-support simulations.

Mud Volcanoes

Mud volcanoes (MVs) are dynamic features of compressional tectonic regimes that occur in both marine ([Napoli et al. 2025](#)) and terrestrial settings ([Milkov 2000](#)). They vent fluids and fine-grained sediments dominated by methane, along with CO₂ and other hydrocarbons ([Mazzini and Etiope 2017](#)), and display diverse morphologies ranging from low mud flows to steep cones ([Kopf 2002](#); [Dimitrov 2002](#)). MVs transport clay-rich sediments, rock fragments, and subsurface microorganisms from the deep subsurface to the surface, providing opportunities to study extremophiles and their metabolic networks in fluid- and mineral-rich habitats ([Ijiri et al., 2018](#); [Lee et al., 2018a](#); [Rajendran et al., 2025](#)). While extensive research on marine MVs revealed necessary syntrophic interaction for microbial survival, between anaerobic methanotrophs (ANMEs) and their sulfate-reducing partners, representative of Early Earth scenarios ([Orphan et al. 2001](#); [Lösekann et al. 2007](#); [Pachiadaki et al.](#)

[2011; Lee et al. 2018b; Omoregie et al. 2008; Ruff et al. 2019; Niemann et al. 2006](#)), the terrestrial MV's have recently gained traction ([Miyake et al. 2023; Merkel et al. 2021; Tu et al. 2017](#)). Terrestrial MV's provide a more accessible systems critical for understanding the limits of Earth's deep biosphere and for interpreting methane emissions and mud-like morphologies on Mars and other worlds (Krýza et al. 2025; Komatsu et al. 2016; Skinner and Mazzini 2009; Hosein et al. 2014).

Prominent terrestrial MVs occur across the Makran accretionary prism, including the Hingol complex, Gwadar, Ormara, and Lasbela regions, where large cones such as Chandragup I (~90 m) episodically discharge methane-rich fluids and fine sediments ([Delisle et al. 2002; Kassi et al. 2014](#)). In the northern Andaman forearc basin, regional compression, overpressured shales, and mass-transport deposits control MV formation ([Ankush et al. 2024; Ray et al. 2013; Kumar et al. 2021](#)). Terrestrial MVs in the Andaman Islands actively emit methane-rich gases, water, and mud breccia, with occasional eruptions triggered by seismic activity ([Chaudhuri et al. 2012](#)). Microbial studies in the Andaman and Nicobar MVs have so far catalogued broad bacterial groups and heterotrophic communities ([Amareesan et al., 2018; Meena et al., 2023; Manna et al., 2021](#)), but targeted enrichment of extremophiles adapted to reduced, methane- and hydrocarbon-rich conditions remains limited.

The Himalayan Range

The Himalayan and adjoining mountain ranges, collectively known as the Third Pole, form the highest terrain on Earth (Altitude: 4 km, Area: ~595000 km²) hosts extreme environments including glaciers, natural caves, geothermal springs, saline ecosystems, and ophiolites (Supplementary Table 3; Figure 3, 4). The origin of the Himalayan range lies in the ongoing collision between the Indian and Eurasian plates, beginning around 50 Ma ago, which uplifted the Tethys Ocean floor and produced the world's tallest mountains. This tectonic setting preserves fragments of ancient oceanic lithosphere in the form of ophiolites, which are emplaced sections of oceanic crust and mantle typically found within suture zones ([Kelemen et al. 2023; Coleman 1977; Dickinson 1971](#)). Ophiolites are frequently composed of peridotite, gabbro, and basalt, displaying evidence of alteration by seawater. Serpentinization of ultramafic rocks in ophiolites, results in generation of hydrogen, methane, and other reduced compounds that provide chemical energy for microbial life ([Klein et al. 2013; McCollom and Bach 2009; Moody 1976](#)). These processes are central to understanding how water–rock interactions sustained early life on Earth ([Schwander et al. 2023](#)) and how biosignatures might be preserved in mineralized fractures, offering analogues for habitable

environments on Mars ([Emran et al. 2025](#); [Schulte et al. 2006](#); [Oze and Sharma 2005](#)) and icy moons such as Europa and Enceladus ([Holm et al. 2015](#); [Taubner et al. 2018](#); [Guo and Eiler 2007](#)). The well-preserved Oman (Semail) Ophiolite remains the most complete section of oceanic lithosphere exposed on land, making it a global benchmark for studying oceanic crust formation, alteration, and serpentinite-hosted microbial ecosystems ([Rempfert et al. 2023](#); [Lima-Zaloumis et al. 2022](#); [Rempfert et al. 2017](#); [Fones et al. 2019](#)). Across South Asia, ophiolite remnants of the Neo-Tethys Ocean occur along the Indus Suture Zone in Ladakh and Pakistan, the Indo–Myanmar ranges, the Indus–Tsangpo belt in Tibet, and the Andaman Islands ([Villalobos-Orchard et al. 2025](#); [Bhat et al. 2022](#); [Ullah et al. 2020](#); [2023](#); [Jalil et al. 2023](#); [Pedersen et al. 2010](#)). Some sites report hydrocarbon-rich fluids with measurable methane ([Das et al. 2017](#); [Sachan et al. 2007](#)). Yet, microbiological investigations across these ophiolites remain sparse, highlighting the need for systematic studies to assess their role as planetary analogues for habitability.

The Himalayan cryosphere contains the largest ice volume outside the polar regions ([Bolch et al., 2012](#)). It hosts valley and cirque glaciers, ice caps, glacial lakes, permafrost zones, and alpine wetlands shaped by interactions among tectonics, topography, and climate. Glacier-fed lakes and streams are significant carbon sources, emitting up to $1746 \pm 139 \text{ mg C m}^{-2} \text{ d}^{-1}$ and $1960 \pm 176 \text{ mg C m}^{-2} \text{ d}^{-1}$, respectively, largely through dissolved inorganic carbon fluxes ([Shukla et al. 2023](#)). These emissions, along with CO_2 degassing from geothermal hot springs, estimated at $1.5 \pm 1.0 \times 10^{11} \text{ mol yr}^{-1}$ ([Pradhan and Sen 2024](#)), derive primarily from deep metamorphic and mantle processes. As glacial lakes continue to expand rapidly, these environments provide valuable sites to study microbial survival, metabolism, and biosignature production in cold, carbon-rich, UV-exposed settings analogous to early Earth and icy planetary bodies. Despite their importance, microbial processes driving carbon cycling in these systems remain poorly characterized. Glaciers, permafrost, and subglacial zones are considered as planetary analogues to icy habitats on Mars, Europa, and Enceladus ([Garcia-Lopez and Cid 2017](#)). Microorganisms in these environments are polyextremophiles, tolerating low temperatures ([Margesin and Miteva, 2011](#); [Cavicchioli, 2016](#)), high radiation ([Zhang et al. 2023](#)), desiccation ([Brenning and Trombotto 2006](#)), nutrient limitation and surviving by metabolizing sulfur, methane, and iron under subglacial conditions independent of sunlight ([Anesio et al. 2017](#); [Bourquin et al. 2022](#)). Himalayan glaciers ([Stres et al. 2013](#); [Venkatachalam et al. 2015](#); [Rafiq et al. 2017](#); [Kumar et al. 2019](#); [Dhakar and Pandey 2020](#); [Singh et](#)

[al. 2024](#)) and rivers ([Paudel Adhikari et al. 2019](#); [Suyal et al. 2022](#)) host diverse psychrophilic microbes, some capable of degrading complex substrates ([Sanyal et al. 2018](#); [Ali et al. 2025](#)), offering insights into biosignature preservation and prebiotic chemistry under cryogenic conditions.

Geothermal systems (hot springs) further expand the Himalayan analogue landscape. The hot springs, characterized by high pH from carbonate-rich fluids, support thermophilic and chemotrophic microbial communities and are considered plausible sites for prebiotic chemistry. On Mars, mineralogical evidence of carbonates, phyllosilicates, and silica suggests that alkaline fluids once existed ([Hurowitz et al. 2023](#)). Terrestrial examples include the Puga ([Roy et al. 2020](#); [Mondal et al. 2022](#)), Chumathang ([Anu et al. 2024](#)), Panamik ([Mondal et al. 2022](#); [Choudhary et al. 2024](#); [Anu et al. 2024](#)) geothermal hot springs, which host diverse phototrophs and chemotrophs, providing an analogue for initial investigations on biosignature studies ([Ansari et al. 2025](#)). Terrestrial hot springs, with wet–dry cycling, are more conducive to prebiotic polymerization than deep-sea vents ([Westall et al. 2018](#); [Des Marais and Walter 2019](#); [Deamer and Georgiou 2015](#)). In addition, underexplored Himalayan alkaline lakes such as Tso Kar, Kyagar Tso, and Tso Moriri provide further opportunities to investigate biosignature preservation in brine-rich systems ([Pandey et al. 2020](#)).

Together, ophiolites, glaciers, permafrost, and geothermal systems of the Himalayas constitute a diverse natural laboratory. These environments allow investigation of water–rock reactions, carbon fluxes, microbial adaptations, and biosignature preservation under extreme conditions, offering analogues highly relevant to early Earth and potentially habitable planetary settings.

The Indian Ocean

The Indian Ocean formed during the breakup of Gondwana in the Mesozoic, with seafloor spreading initiating along the Carlsberg and Central Indian ridges around 140–120 Ma ([Gaina et al. 2007](#); [2015](#)), hosts many deep sea vents and trenches for exploration (Supplementary Table 4; Figure 3). Its rock record preserves key tectonic and magmatic features, including slow- to ultra-slow spreading ridges, fracture zones, seamount chains, and thick sedimentary sequences influenced by the Himalaya and monsoonal systems. Together, these archives document the evolution of oceanic lithosphere, plate reorganization, and paleoceanographic change across the Cenozoic. Within the Indian Ocean,

oxygen-depleted habitats, including oxygen minimum zones, hadal trenches, hydrothermal systems, and cold seeps, provide valuable analogues for extraterrestrial environments.

Deep sea hydrothermal vents found along mid-ocean ridges, where tectonic plates diverge, allow seawater to circulate through fractured crust and interact with underlying magma. These environments replicate extreme conditions expected on icy moons such as Europa and Enceladus, including high pressure, low temperature, absence of sunlight, and chemical energy derived from water–rock interactions ([McClain et al. 2022](#); [Aguzzi et al. 2024](#)). They serve as natural laboratories for evaluating biosignature production and strategies for life detection in extraterrestrial oceans ([George 2020](#); [German et al. 2022](#)) as they sustain microbial communities reliant on chemosynthesis rather than photosynthesis ([Ricci and Greening 2024](#); [Sogin et al. 2020](#); [Dick 2019b](#)). Evidence from the Galileo mission ([Zimmer et al. 2000](#)) and the Cassini mission ([Barge and Rodriguez 2021](#)), supported by subsequent studies ([Barge and White 2017](#); [Glass et al. 2022](#)), indicates that hydrothermal activity may occur within the subsurface oceans of Europa and Enceladus, creating conditions capable of supporting methanogenesis ([Holden and Sistu 2023](#); [Lyu et al. 2018](#)), while the ongoing Europa Clipper mission seeks to assess Europa’s subsurface ocean directly ([Roberts et al. 2023](#)). The Indian Ocean Ridge System, subdivided into the Central (CIR), Southwest (SWIR), and Southeast Indian Ridge (SEIR), is a slow- to ultra-slow spreading system hosting diverse vent fields. Since their first discovery in 2001 ([Van Dover et al. 2001](#)), notable sites include Kairei, Edmond ([Van Dover et al. 2001](#)), Solitaire, Dodo ([Nakamura et al. 2012](#)), and Onnuri ([Lim et al. 2022](#)) on the CIR; Mount Jourdanne ([Münch et al. 2001](#)), Longqi ([Tao et al. 2012](#)), Duanqiao ([Yang et al. 2017](#)), Tiancheng ([Zhou et al. 2018](#)), Old City ([Lecoeuvre et al. 2021](#)), and Carlsberg ridge ([Qiu et al. 2021](#); [Liang et al. 2023](#); [Cai et al. 2024](#)) on the SWIR; and Pelagia ([Han et al. 2018](#)) on the SEIR. These fields exhibit variable geology, from basaltic to ultramafic hosts, supporting rich mineral formations including massive sulfides, iron-oxyhydroxides, barite, and talc ([Perez et al. 2021](#); [van der Most et al. 2023](#); [Thomas et al. 2024](#); [Ta et al. 2024](#)). Microbial studies using isolations, 16S rRNA sequencing, and metagenomics reveal assemblages adapted to vent-specific gradients in fluids, chimneys, and plumes ([Li et al. 2016](#); [Huang et al. 2023](#); [Adam-Beyer et al. 2023](#); [Namirimu et al. 2022](#); [Wee et al. 2021](#); [Bai et al. 2021](#); [Han et al. 2018](#); [Ding et al. 2017](#); [La Duc et al. 2007](#)). Thermophiles and chemolithoautotrophs drive sulfur, iron, and methane cycling ([Zhong et al. 2022](#); [Cao et al. 2014](#); [Surya Prakash et al. 2025](#); [Wee et al. 2021](#)), while symbiotic communities support endemic fauna

such as Rimicaris shrimp and Scaly-foot snails ([Bai et al. 2021](#); [Methou et al. 2022](#)). These ecosystems highlight metabolic diversity relevant to potential chemolithotrophic life on icy moons.

Deep-sea trenches develop at convergent margins where one tectonic plate is subducted beneath another, creating the deepest depressions on Earth's seafloor. Hadal trenches ([Jamieson et al. 2010](#); [Jamieson 2020](#)), at depths greater than 6,000 meters, offer additional insights into microbial survival under extreme hydrostatic pressure, low temperature, and limited nutrients ([Tyler 2003](#); [Jørgensen and Boetius 2007](#)). Redox stratification within trench sediments produces distinct oxic, nitrogenous, and ferruginous zones ([Luo et al. 2018](#); [Schauberger et al. 2021](#)), and pressure effects may even induce anaerobic metabolism in otherwise oxic conditions ([Yang et al. 2024](#)). Until today, the Mariana Trench remains the most studied hadal site, with work on geomorphology, geochemistry, microbial adaptations, and autonomous sampling technologies revealing that nearly 89% of recovered microbial taxa remain unclassified ([Dietrich et al. 1978](#); [Kato et al. 1998](#); [Tarn et al. 2016](#); [Nunoura et al. 2018](#); [Liu et al. 2019](#); [Li et al. 2021](#); [Xiao et al. 2025](#)). Other key trenches include the Japan, Izu–Ogasawara ([Hiraoka et al. 2020](#); [Arakawa et al. 2005](#)), and Kermadec systems ([Zhang et al. 2024](#)). Recent discoveries in the Kuril–Kamchatka and Aleutian Trenches revealed chemosynthetic communities spanning 2,500 km at depths exceeding 9,000 m ([Peng et al. 2025](#)). The Indian Ocean hosts the Java Trench, that reaches hadal depths and has remained largely unsampled for decades ([Jamieson 2020](#)). A 2019 expedition using a full-depth submersible revealed diverse habitats and microbial-associated features, including putative chemolithoautotrophic bacterial mats, underscoring the urgent need for targeted exploration of this trench to understand extreme microbial ecosystems better ([Jamieson et al. 2022](#)).

To maximize the astrobiological relevance of deep-sea research, Earth-based studies must be explicitly linked with models of extraterrestrial hydrothermal environments. Laboratory pressure reactors simulating icy moon conditions allow controlled investigations of microbial survival, metabolism, and biosignature production. Observations of plume chemistry and dynamics at Earth's vents may also refine strategies for remote sensing of plumes on icy moons. Expanded exploration of Indian Ocean environments with autonomous submersibles and robotic platforms will not only deepen understanding of microbial survival strategies and biogeochemical processes but also provide critical test beds for technologies being developed for space missions ([Feng et al., 2022](#); [O'Neill, 2021](#); [Liang et al., 2021](#)).

Extreme Environments as Planetary Analogues for Habitability

Life, as currently understood, requires fundamental physicochemical conditions that support processes such as metabolism, growth, and reproduction ([Falkowski et al. 2008](#)). These requirements are largely derived from Earth, where microbial life represents the earliest and most adaptable form of biology ([Nealson and Conrad 1999](#)). Although extraterrestrial life may differ in its chemistry, it is still expected to follow universal principles ([Benner 2010](#); [Cockell et al. 2016](#); [Hoehler 2007a](#); [McKay 2022](#); [Méndez et al. 2021](#)). The most critical requirement is the presence of a solvent. On Earth, the unique properties of water, including polarity, heat capacity, and its ability to dissolve diverse compounds, make it indispensable for sustaining life. Therefore, planetary bodies and exoplanets that host liquid water, even transiently, are considered prime targets for exoplanet research ([Kite and Ford 2018](#); [Goldblatt 2015](#)). Another essential requirement is the availability of key elements such as carbon, hydrogen, nitrogen, oxygen, phosphorus, and sulfur (CHNOPS), which form the backbone of biomolecules ([Rossetto and Mansy 2022](#)). Microorganisms derive energy and assimilate essential elements for growth and reproduction from their geochemical environment through redox-active compounds that exist in chemical disequilibrium ([Hoehler 2007b](#)). This concept holds particular significance in astrobiology, as it implies that planets or moons exhibiting active water–rock interactions may provide the chemical energy necessary to sustain life in environments devoid of sunlight.

Earth’s extreme environments illustrate these possibilities. Microorganisms in deserts, deep crustal rocks, hypersaline lakes, and hydrothermal systems (Figure 3) exploit chemical energy from compounds such as iron, sulfur, or hydrogen. Their ability to persist under conditions of high radiation, desiccation, or pressure demonstrates the plausibility of life in dark subsurface habitats on early Earth, Mars, Icy worlds (Europa, or Enceladus) (Figure 4). These adaptations expand our understanding of habitability and provide models for biosignature preservation in energy-limited settings. Beyond informing the search for life, microbial strategies are increasingly relevant to space exploration itself. Insights from analogue studies guide planetary protection policies and inform the design of life-detection instruments. They also highlight how biology might contribute to sustaining long-duration missions. Microorganisms capable of extracting metals from regolith, producing oxygen, or recycling resources could support in-situ resource utilization (ISRU) on the Moon, Mars, and beyond ([Santomartino et al. 2023](#); [Koehler et al. 2023](#); [Cockell et al. 2024](#)). Testing these processes under planetary analogue conditions on Earth enables the development of sustainable and adaptable bioengineering strategies for off-world survival. Integrating such biology–geology systems

into exploration frameworks is therefore critical for both advancing astrobiology and meeting the practical challenges of human spaceflight.

Scientific Gaps from the Indian Astrobiology Perspective

From an astrobiological standpoint, extreme environments across the Indian subcontinent and the Indian Ocean remain underexplored despite their relevance as planetary analogues (Figure 2). Microbial life has been reported from glaciers, geothermal springs, alkaline lakes, hydrothermal vents, permafrost zones, and impact structures such as Lonar crater, yet systematic insights into their ecological roles and astrobiological significance are still limited. In many of these habitats, the physiology and metabolic strategies of extremophiles capable of withstanding high pressure, temperature, salinity, or acidity remain poorly characterized. Subsurface environments, including permafrost layers, subglacial sediments, and deep-sea trenches, are particularly neglected, even though they closely resemble shielded niches that could exist on Mars, Europa, or other planetary bodies.

Key research gaps arise from methodological limitations. Most studies rely on isolated sampling campaigns and culture-based approaches, offering only a partial view of microbial diversity. The limited use of advanced molecular techniques such as metagenomics, metatranscriptomics, and metabolomics restricts our understanding of functional capabilities and ecological dynamics ([Clark et al., 2023](#); [Jansson and Baker, 2016](#); [Krassowski et al., 2020](#)). Without long-term monitoring, the response of microbial communities to seasonal or episodic environmental fluctuations also remains unresolved ([Nguyen et al. 2021](#); [Gunnigle et al. 2017](#); [Whitaker and Banfield 2005](#)). Another critical gap lies in the study of biosignature preservation. While pigments, lipids, isotopic signals, and biominerals provide promising targets, little is known about how these markers persist under the unique geochemical and physical conditions of Indian extreme environments ([Barge et al. 2022](#); [Campbell et al. 2015](#); [Hays et al. 2017](#); [Moore et al. 2022](#); [Summons et al. 2011](#)). Addressing these gaps is essential for developing effective life-detection strategies for future planetary missions.

Preservation of these geological sites is equally critical. Many analogue sites, in and around Indian subcontinent and the Indian ocean, are increasingly threatened by unregulated geotourism, microbial contamination, and anthropogenic alteration of fragile mineralogical and geochemical gradients ([Dowling 2010](#); [Santos and Brilha 2023](#)). Extreme environments, as sensitive indicators of

environmental change, respond rapidly to shifts in temperature, precipitation, and chemical fluxes. Protecting their integrity is therefore essential not only for advancing astrobiology but also for sustaining long-term ecological and climate monitoring. Linking planetary exploration with Earth's analogue sites underscores a dual responsibility: advancing scientific discovery while safeguarding irreplaceable natural heritage from irreversible human impact.

Realizing the full potential of these environments requires coordinated interdisciplinary research that integrates geomicrobiology, geochemistry, planetary science, and remote sensing. Establishing long-term observatories, deploying in situ analytical tools, and fostering regional and international collaborations will be vital for building a robust framework for analogue research. Beyond their astrobiological value, these sites serve as natural laboratories for studying early Earth conditions and biogeochemical cycling of carbon, sulfur, and methane under natural constraints, thereby informing both planetary exploration and climate science.

Challenges in Studying Analog Sites hosted in the Indian subcontinent

Astrobiologically motivated geomicrobiological research in planetary analogue sites faces several challenges that affect both sampling and data interpretation. Foremost are accessibility and logistical constraints. High-altitude regions such as Ladakh and remote geothermal zones in the Himalayas often require specialized transport, acclimatization, and multiple administrative permissions. These limitations reduce the frequency and duration of field campaigns, restricting the resolution of temporal and spatial microbial assessments. Environmental heterogeneity and seasonal variability further complicate the identification of representative sampling windows. A hybrid strategy that combines targeted seasonal expeditions with autonomous in situ monitoring could help overcome these limitations. Portable analytical tools, remote sensing platforms, and sentinel sensor networks can extend temporal coverage, while standardized low-impact sampling protocols preserve site integrity and ensure reproducibility. Long-term collaborations with local institutions and authorities are also essential to streamline permitting and enable sustained access.

Addressing these concerns requires strict contamination control, including not only sterile or UV-sterilized tools but also comprehensive use of field blanks and negative controls. Establishing restricted-access research zones can further reduce human disturbance, while multi-seasonal monitoring of physicochemical parameters (pH, salinity, redox, mineral composition) coupled with high-resolution molecular profiling (e.g., shotgun metagenomics, amplicon sequencing) will help establish reliable baselines. Incorporating computational decontamination pipelines ([Murray et al.](#)

[2015; Schmieder and Edwards 2011; Lu and Salzberg 2018](#)) can then separate indigenous microbial signals from transient contaminants. Closely related to this challenge is the lack of standardized procedures for aseptic sampling, in situ preservation, and metadata reporting, which limits reproducibility and cross-site comparisons. Developing unified protocols aligned with international standards such as MIXS ([Field et al. 2011](#)) and Darwin Core ([Wieczorek et al. 2012](#)) is therefore critical. This can be achieved through interdisciplinary workshops that bring together planetary scientists, geomicrobiologists, and geochemists to co-design context-specific methods, followed by pilot testing across diverse analogue sites. Publishing the resulting framework as an open-access, modular protocol, with training modules and field-ready checklists, would promote consistency, enhance data comparability, and support long-term multi-site analysis.

Finally, the establishment of dedicated infrastructure near key analogue sites, such as field stations, environmental monitoring platforms, and sample repositories, would strengthen the research ecosystem. Policy support is needed to designate protected analogue research zones, ensuring both long-term access and preservation of site integrity. Together, these measures would allow planetary analogue environments to be studied with the rigor required for astrobiological relevance, positioning them as critical contributors to global efforts in understanding habitability, biosignature preservation, and planetary exploration.

Future perspectives

The Indian subcontinent and its adjoining oceans host an exceptional diversity of extreme environments, ranging from high-altitude cold deserts and geothermal springs in the Himalayas to hypersaline lagoons, deep caves, submarine hydrothermal systems, and cold seeps along the Indian Ocean ridges. These settings serve as valuable planetary analogues, offering insights into habitability and biosignature preservation beyond Earth. To harness this potential, the Indian Space Research Organisation (ISRO), alongside global space agencies, should prioritize the systematic identification and designation of analogue sites across the region.

A key step forward would be the establishment of dedicated planetary analogue observatories at strategically selected field sites. These observatories could combine in situ measurements, autonomous sensors, and remote sensing data to generate multi-scale datasets that capture both short-term variability and long-term climatic and geological trends. Emerging technologies, including portable genomic sequencers, autonomous loggers, underwater drones, and satellite-linked sensor networks, can help overcome logistical constraints and expand monitoring capacity. Integrating

573 ISRO's Earth observation capabilities with international datasets would further enable high-
574 resolution mapping of extreme habitats and microbial community dynamics. Fostering
575 interdisciplinary investigations that connect microbiologists, planetary scientists, oceanographers,
576 and geochemists will help cultivate the expertise necessary for research across varied environments.
577 Moreover, collaborative frameworks among universities, research institutes, government agencies,
578 and space organizations will streamline fieldsite access, share infrastructure, and coordinate long-
579 term field campaigns. Embedding these efforts within open-access data-sharing platforms will
580 amplify the international visibility of Indian research while promoting responsible stewardship of
581 fragile terrestrial and marine analogue sites.

582 In the long term, the Indian subcontinent and Indian Ocean would emerge as a global hub for
583 planetary analogue science, directly supporting exploration of the Moon, Mars, and icy worlds. These
584 environments provide natural laboratories for testing life-detection technologies, validating sample
585 handling protocols, and refining hypotheses about life's adaptability under extreme conditions. By
586 aligning national priorities with international astrobiology roadmaps and coupling exploration with
587 conservation, India will be well-positioned to make a distinctive contribution to the search for life
588 beyond Earth while safeguarding its own unique analogue environments.

589 **Author Contributions**

590 Conceptualization: Y.J.; Methodology: Y.J. and S.D.; Validation: Y.J. and S.D.; Formal analysis:
591 Y.J. and S.D.; Resources: Y.J.; Data curation: Y.J. and S.D.; Writing (original draft, review, and
592 editing): Y.J.; Visualization: S.D.; Supervision: Y.J.

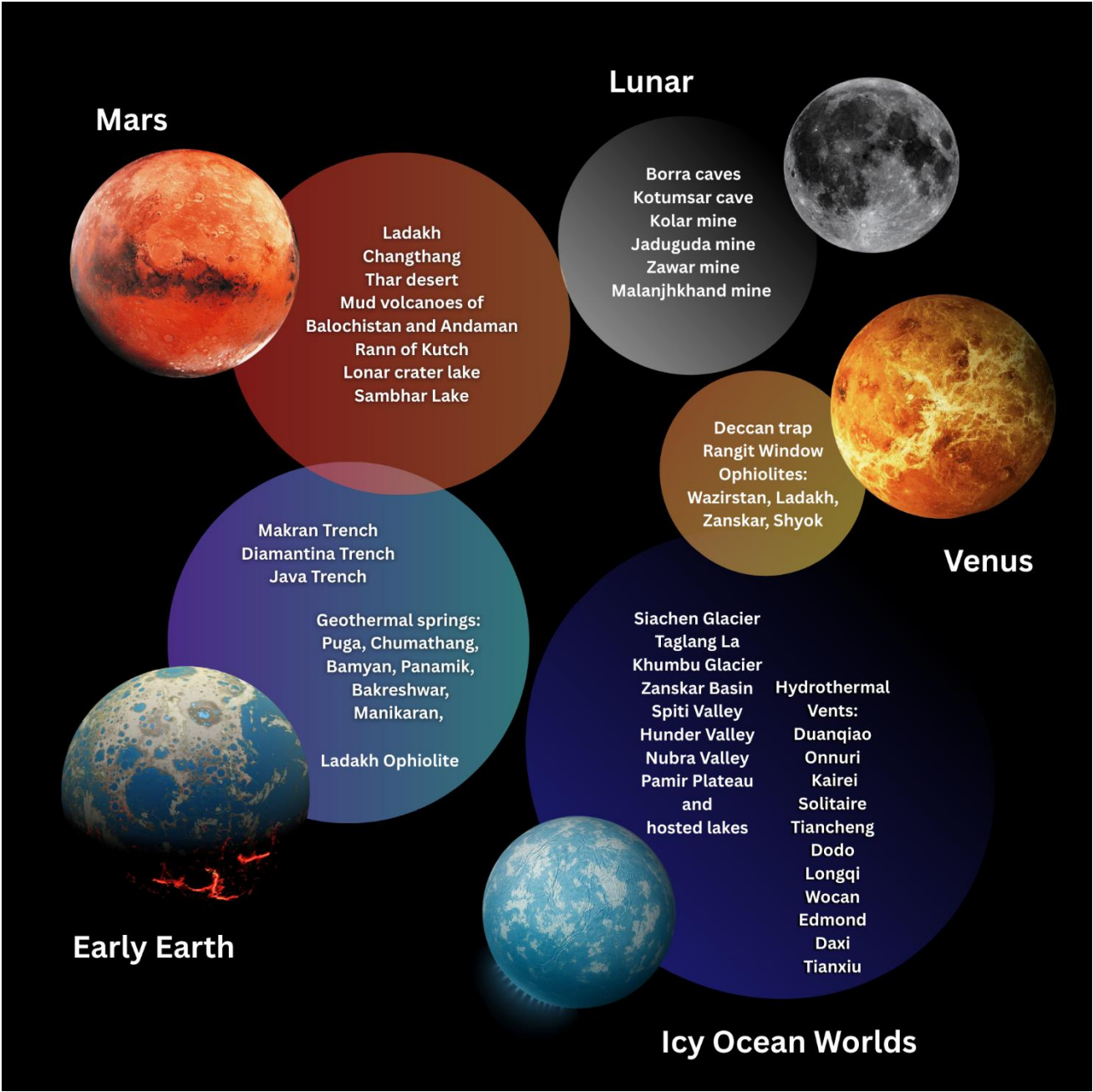
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602 astrobiological framework within the Indian space research community.

605 **Figure 1:** Planetary analog environments across the Indian subcontinent and the Indian ocean basins
606 that represent planetary bodies of astrobiological interest.



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Figure 2. Schematic representation of extreme environments across the Indian subcontinent and the Indian Ocean that serve as planetary analogues. Terrestrial ecosystems include mud volcanoes, saline habitats, desert sandstorms, glaciers, ophiolites, geothermal springs, high-altitude lakes, and natural caves and mines. Marine ecosystems of the Indian Ocean remain relatively underexplored, with documented sites including deep-sea hydrothermal vents and subsurface trenches. This figure is adapted and inspired by Merino et al. (2019).

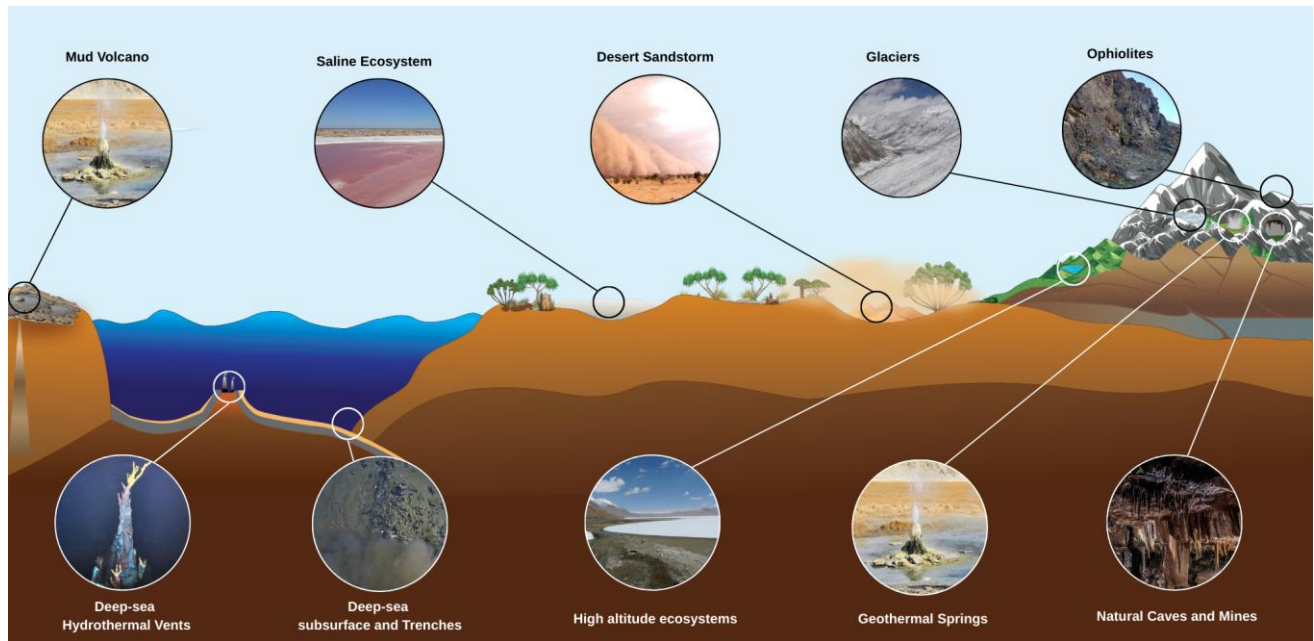


Figure 3: Geographic distribution of extreme environments, that can serve as planetary analogues, overlaid on a hypsometric map of the Indian subcontinent and adjoining Indian Ocean. The left panel illustrates the regional distribution of terrestrial and marine sites across diverse ecosystems, while the right panel provides a detailed view of the Himalaya and Tibetan Plateau, highlighting clusters of high-altitude and glacial environments. The underlying topography and bathymetry emphasize the environmental gradients that shape the diversity of analogue sites. Geological base map compiled from GEBCO Compilation Group (2025).

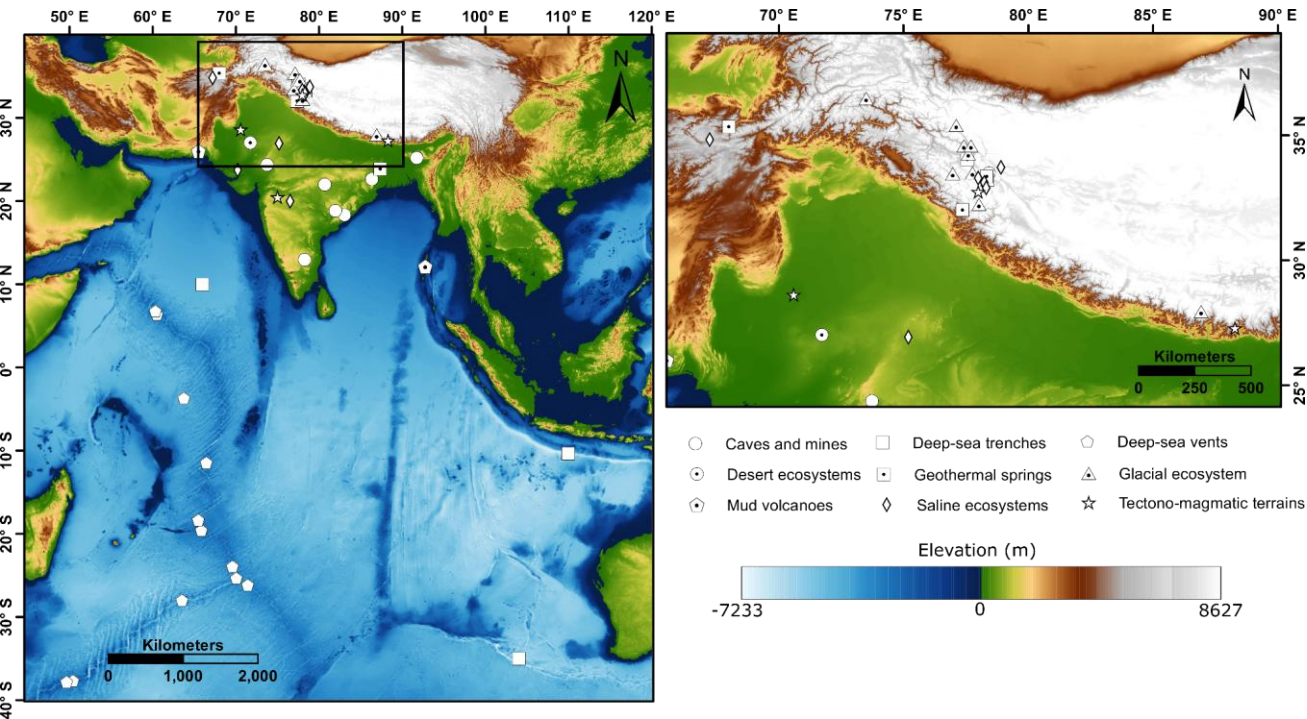
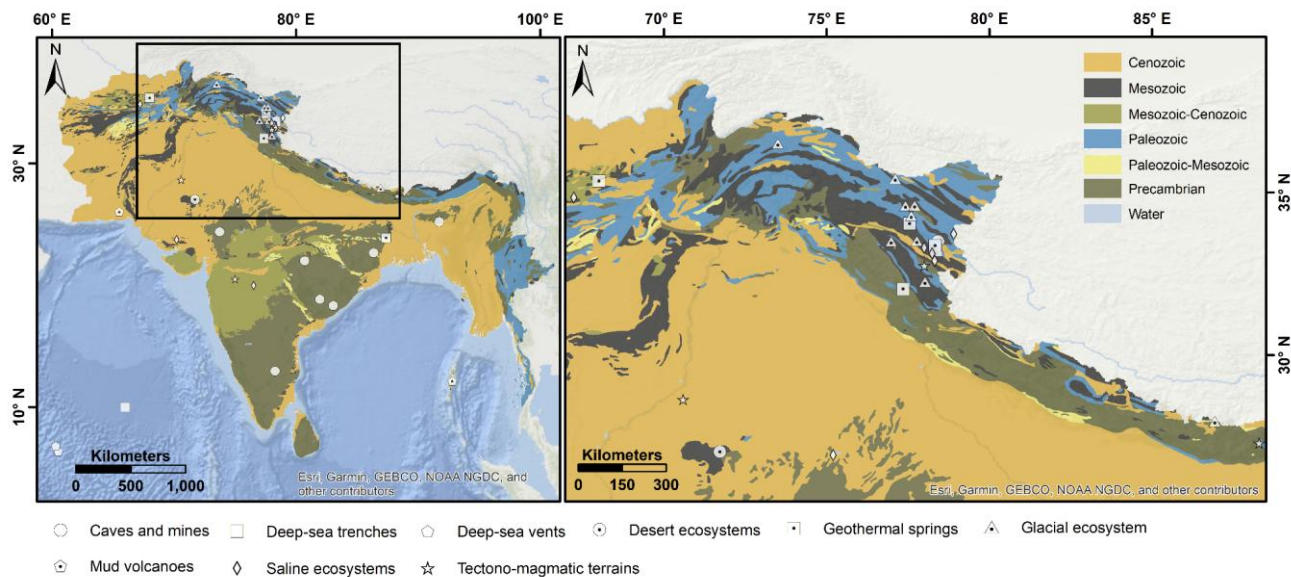


Figure 4: Geographic distribution of extreme environments, that can serve as planetary analogues, overlaid on a generalized geological map of the bed rock of the Indian subcontinent. The underlying geological framework (Cenozoic, Mesozoic, Paleozoic, and Precambrian units) highlights the lithological diversity associated with these analogue sites. Geological base map compiled from Wandrey, C.J., 1998, *Geologic map of South Asia (geodag)*.



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Supplementary Material

Table 1: Astrobiology field sites around the globe that serve as planetary analogues.

Field site	Analog For	Key Features	Latitude	Longitude	Reference
Dallol Geothermal Area, Ethiopia	Early Earth, Mars	Acidic pools, high salinity, hydrothermal	14.20	40.30	Cavalazzi et al. 2019
Gakkel Ridge, Central Arctic Ocean	Europa and Enceladus	Hydrothermal Vent Fields	81.00	5.00	Ramirez-Llodra et al. 2023
Lost City, Mid-Atlantic Ridge, Atlantic Ocean	Europa and Enceladus	Hydrothermal Vent Fields	30.13	-42.12	Amador et al. 2017
Juan de Fuca, Pacific Ocean	Europa and Enceladus	Hydrothermal Vent Fields	48.26	-123.36	Kelley et al. 2012; Osinski et al. 2006
Mariana trench, Pacific Ocean	Europa and Enceladus	Hadal trench	11.35	142.20	Preston and Dartnell 2014
Lake Vostok, Antarctica	Europa and Icy Moons	Subglacial lake, high pressure, darkness	-77.00	106.00	Cassaro et al. 2021
Kīlauea Volcano, Hawaii	Lunar and Mars	Volcanism, basaltic terrain	19.43	-155.29	Bergsten et al. 2021; D’Incecco et al. 2024; Seelos et al. 2010
Haughton Impact Crater, Canada	Lunar and Mars	Permafrost-preserved impact crater	75.00	-89.50	Spray et al. 2010; Cloutis et al. 2015
Atacama Desert, Chile	Mars	Hyperarid, high UV, Mars-like soil	-24.00	-69.25	Azua-Bustos et al. 2022; Lebre et al. 2017; Visscher et al. 2020
Death Valley, USA	Mars	Extreme heat, dryness, desiccation	36.50	-117.00	Farr 2004; Douglas et al. 2008
Icelandic Volcanic Fields	Mars	Basaltic geology, subglacial hydrothermalism	65.00	-17.00	Voigt et al. 2025; Fagents and Thordarson 2007
Rio Tinto, Spain	Mars	Acidic, iron-rich river, extremophiles	37.68	-6.57	Amils et al. 2014; 2007; Gómez et al. 2011
Tibetan Plateau, China	Mars	Cold, arid, high-UV, hypoxic	32.00	91.00	L. Liu et al. 2022; Sun et al. 2018
Strýtan Hydrothermal Field, Iceland	Mars and Europa	Alkaline Hydrothermal vent	65.82	-18.12	Twing et al. 2022
McMurdo Dry Valleys, Antarctica	Mars and Icy Moons	Cold, dry, permafrost, low nutrients	-77.50	161.50	Salvatore and Levy 2021
Svalbard Archipelago, Norway	Mars, icy moons	Glacial and permafrost environment	78.20	15.60	Kołodziejczyk et al. 2025; Hauber et al. 2011

Table 2: Compilation of putative planetary analogue sites across the Indian subcontinent classified by ecosystem type, geological context, mineralogy, and planetary analog relevance. Each site is characterized by its host lithology, geologic age, geochemical parameters (pH, temperature, salinity), and associated minerals. Salinity values are standardized as: Low (<1,000 mg/L TDS), Mid (1,000–10,000 mg/L TDS), and High (>10,000 mg/L TDS). The table highlights analog links to planetary bodies (Mars, Moon, Venus, and Icy Ocean worlds such as Europa and Enceladus) and to Early Earth. References provided for each field site correspond to published geological and geomicrobiological studies.

Environment	Field site	Lithology (hosted rocks)	Geological age of host rock (Ma)	Associated Minerals	Lat.	Long.	Altitude (m)	pH	Temp (°C)	Salinity	Planetary Analog	Rationale	References
Caves and mines	Borra Caves	Terrestrial subsurface karst cave hosted in proterozoic (Mesoproterozoic–Neoproterozoic) limestones of the Cuddapah Supergroup	5700	Calcite, aragonite, speleothems, oxides	18.33	83.05	705				Lunar/Mars	Lava tubes and fractured crust on Mars/Moon provide radiation-shielded habitats analogous to Earth's subsurface caves and mines.	Baskar and Baskar 2022; Baskar et al. 2009; 2011
	Kotumsar Cave	Terrestrial subsurface karst cave hosted in Proterozoic (Mesoproterozoic–Neoproterozoic) limestones of the Chhattisgarh Supergroup	1000	Calcite, aragonite, speleothems, oxides	18.85	81.95	560						
	Kolar Gold Fields	Terrestrial subsurface mine hosted in Archaean (Late Archaean, ~2.6 Ga) greenstone belt volcanics and granitoids	2500	Gold, quartz, sulfides (pyrite, arsenopyrite)	12.95	78.27	900						Siddaiah and Rajamani 1989; Reddy et al. 2017
	Jaduguda Uranium Mine	Terrestrial subsurface mine hosted in Proterozoic (Paleoproterozoic, ~1.6–1.4 Ga) metasediments and granites	1400	Uraninite, chalcopryrite, pyrite, chlorite, quartz	22.65	86.35	400						Pal et al. 2011
	Zawar Mines	Terrestrial subsurface mine hosted in Proterozoic (Paleoproterozoic, ~1.8–1.6 Ga) carbonates and volcanics of the Aravalli Supergroup	1600	Lead, zinc (sphalerite, galena), pyrite, carbonates	24.37	73.73	550						Pandey et al. 2007; Equeenuddin et al. 2017; Sikka 1989
	Malanjkhand Copper Mine	Terrestrial subsurface mine hosted in Proterozoic (Late Archaean–Early Proterozoic, ~2.5–2.0 Ga) granitoid-hosted copper deposit	2000	Chalcopryrite, bornite, quartz, magnetite, pyrite	22.00	80.70	600						
Desert ecosystems	Thar Desert	Proterozoic Marwar Supergroup, along with isolated outcrops of the Malani Igneous Suite, the Erinpura Granites, and the Sirohi Group. In the southwest (Jaisalmer and Barmer Basins), Mesozoic and Tertiary sedimentary deposits are found, bordered by the extensive Quaternary alluvium.	600	Dune sands chiefly quartz; playas dominated by halite with gypsum/anhydrite, calcite/dolomite and detrital quartz–feldspar–mica	27.02	71.72	200		-5 to 50		Mars	Hyperarid conditions and sandstorms serve as analogs for the Martian environment, where extreme aridity, dust storms, and limited water availability shape surface processes and habitability.	Khormali and Monger 2020; Sarkar et al. 2019

Environment	Field site	Lithology (hosted rocks)	Geological age of host rock (Ma)	Associated Minerals	Lat.	Long.	Altitude (m)	pH	Temp (°C)	Salinity	Planetary Analog	Rationale	References
Geothermal springs	Bakreswar	Springs emerge along fractures associated with Precambrian granitic rocks and sedimentary remnants of the Gondwana sequence (ranging from the Lower Permian to the Middle Jurassic)	100	Primarily granite within the Archean Chotanagpur Gneissic Complex	23.88	87.37	79	7-8	50 to 70	Low	Early Earth	Hydrothermal springs were likely sites for prebiotic chemistry and early life emergence.	Das et al. 2022; Majumdar et al, 2000; Shukla et al, 2022;
Mud volcanoes	Balochistan	Modern (Holocene) geothermal systems in Makran accretionary prism (mainly Oligocene–Miocene flysch	100	Silica sinter (opal-A), calcite/aragonite; sandstones/shales rich in quartz, feldspar, clay minerals.	26.01	65.51	200	6-7	30 to 90	Moderate	Mars	Observed mud volcano-like structures on Mars suggest cryovolcanic or sedimentary-fluid extrusion processes analogous to Earth mud volcanoes.	Delisle et al, 2002; Kassi et al., 2014; Yaseen et al, 2021
	Andaman Islands	MVs are hosted by the Eocene trench-slope deposits, called the Mithakhari Group, the clasts in the mud breccia are primarily from this lithology. Clasts of basalt belonging to the ~94 Ma old Ophiolite Group	94	illite, kaolinite, volcanic glass, brine and sulfide nodules	12.13	92.79				High			Chaudhuri et al. 2012; Achyuthan and Eastoe, 1999
Saline ecosystems	Rann of Kutch	Holocene supratidal salt flats (sabkhas) with annual flooding/evaporation	66	Surface halite crusts; gypsum grows within clays/sands; brines enriched in boron.	23.73	70.22	15	8-10	-2 to 48	High	Mars	Evaporitic basins mineral deposits record past aqueous activity and possible habitats for microbial life.	Biswas et al., 2021; Chavan et al, 2022; Singh et al.,2021
	Lonar Crater Lake	Pleistocene impact (~0.57 Ma) excavated into Deccan Traps basalts (~66 Ma)	66	Basalt dominated by plagioclase (labradorite) and pyroxene (augite/pigeonite); shocked glass and maskelynite reported	19.98	76.51	563	10.5	15 to 45	High			Fredriksson et al., 1973; Maloof et al., 2010; Antony et al., 2013
	Sambhar Lake	Holocene–Late Quaternary playa/lacustrine sediments surrounded by proterozoic Aravalli hills	200	Halite, gypsum/anhydrite, lacustrine carbonates; clastic quartz–feldspar–mica influx from catchment	26.92	75.20	360	9-10	7 to 37	High			Sinha et al., 2004; Singh et al., 2024; Sahay et al., 2012, Sarkar et al., 2023; Pal et al., 2020
Tectono-magmatic terrains	Deccan Traps	Late Cretaceous–Paleogene continental flood basalts (~66–64 Ma)	70	Tholeiitic basalts rich in plagioclase, clinopyroxene (augite), ± olivine; secondary zeolites common	20.50	75.00	600		-10 to 45		Early Earth/Venus	Volcanic provinces, oceanic crustal rocks, and ductile crustal processes were widespread on the early Earth and Venus	Sen et al. 2009; Beane et al. 1986; Dutta et al. 2019; 2018; Jungbluth et al. 2016
	Waziristan	Late Jurassic to Early Cretaceous ophiolites	150	braunite, cryptocrystalline quartz, minor hematite	28.63	70.59	100		-5 to 50				Jalil et al., 2023

Table 3: Compilation of putative planetary analogue sites across the Himalayan region classified by ecosystem type, geological context, mineralogy, and planetary analog relevance. Each site is characterized by its host lithology, geologic age, geochemical parameters (pH, temperature, salinity), and associated minerals. Salinity values are standardized as: Low (<1,000 mg/L TDS), Mid (1,000–10,000 mg/L TDS), and High (>10,000 mg/L TDS). The table highlights analog links to planetary bodies (Mars, Moon, Venus, and Icy Ocean worlds such as Europa and Enceladus) and to Early Earth. References provided for each field site correspond to published geological and geomicrobiological studies.

Environment	Field site	Lithology (hosted rocks)	Geological age of host rock (Ma)	Associated Minerals	Lat	Long	Altitude (m)	pH	Temp (°C)	Salinity	Planetary Analog	Rationale	References
Caves and mines	Krem Phyllut	Terrestrial subsurface karst cave hosted in Late Cretaceous–Paleogene limestones of the Shillong Plateau	66	Calcite, aragonite, speleothems	25.25	91.73	1300				Lunar/Mars	Lava tubes and fractured crust on Mars/Moon provide radiation-shielded habitats analogous to Earth's subsurface caves and mines.	Baskar and Baskar 2022; Baskar et al. 2009; 2011
	Mawsmi Caves	Terrestrial subsurface karst cave hosted in Late Cretaceous–Paleogene limestones of the Shillong Plateau	66	Calcite, aragonite, speleothems, silica	25.20	91.73	1200						
Desert ecosystems	Changthang	Quaternary alluvial–lacustrine cover over Ladakh Batholith (Cretaceous–Paleogene granitoids) and Tethyan sedimentary sequence (Paleozoic–Mesozoic carbonates/siliciclastics)	50	Batholith felsics (quartz–K-feldspar–plagioclase–biotite); Tethyan units dominated by calcite/dolomite, quartz arenites and shales (clays).	33.50	78.42	4500		-40 to 25			Hyperarid conditions, high UV, serve as terrestrial analogs for surface environments on Mars and the Lunar regolith, where desiccation, intense radiation, and limited shielding define habitability constraints	Ali et al, 2025
Geothermal springs	Puga	Active geothermal field depositing modern (Holocene) travertine/sinter on Trans-Himalayan basement; quartzo feldpathic gneiss of Late Proterozoic cambrian age at the deepest levels	100	Borates (borax/ulexite), thenardite (Na_2SO_4), halite, opaline silica, plus native sulfur and borax observed	33.22	78.36	4400	7.5-9	30 to 84	Low	Early Earth	Hydrothermal springs were likely sites for prebiotic chemistry and early life emergence.	Gopal et al, 2002, Sarkar et al, 2022; Harinarayana et al, 2006; Ansari et al., 2025; Kumari et al., 2024; Verma et al., 2022
	Chumathang	Holocene hot-spring precipitates (travertine) over Trans-Himalayan host rocks	100	Calcite/aragonite travertine; regional reports of borax/sulfur	33.36	78.33	3960	8-9	40 to 86	Mid			Ansari et al., 2023; Ansari et al., 2025; Verma et al., 2022
	Bamyan	Modern (Holocene) hot-spring activity associated with Quaternary tectonics; hosted in Paleozoic–Mesozoic carbonates and granites.	100	Travertine (calcite), opaline silica, and host-rock minerals (quartz–feldspars, calcite/dolomite from limestones).	35.35	68.00	2500	6-8	20 to 30	Mid			Jawadi et al., 2021
	Panamik	Ongoing Holocene hot-spring deposition in Nubra Valley	60	Spring precipitates include travertine (calcite/aragonite)	34.04	77.55	3200	7-8	60 to 70	Low			Ansari et al., 2023; Ansari et al., 2025; Verma et al., 2022
	Manikaran	Holocene hydrothermal deposits within Higher Himalayan metasedimentary/granitic rocks	0.01	Active travertine (CaCO_3) formation at vents	32.03	77.35	1760	6-7	35 to 94 (spring water)	Low			Husain et al., 2020

Environment	Field site	Lithology (hosted rocks)	Geological age of host rock (Ma)	Associated Minerals	Lat	Long	Altitude (m)	pH	Temp (°C)	Salinity	Planetary Analog	Rationale	References
Glacial ecosystem	Khardung La	Pass cuts Khardung Volcanics (Cenozoic rhyolitic to intermediate lavas) over Ladakh Batholith (Cretaceous–Paleogene)	100	Quartz–feldspar-rich rhyolites/ignimbrites; batholith granodiorite–granite with quartz–K-feldspar–plagioclase–biotite	34.28	77.60	5602		-30 to 15		Icy Ocean Worlds	Mimics outer ice shells of Europa, Enceladus, and other icy worlds	Lakhan et al, 2019; Srikantia and Razdan, 1980
	Siachen Glacier	Quaternary glacier on Karakoram bedrock (High-grade metamorphic complexes and granites)	1.5	Quartz–feldspar–mica gneisses/schists; granitoids with quartz–feldspars–biotite; local amphibolites	35.42	77.11	5400		-50 to 10				Saktura et al, 2023
	Taglang La	Metamorphosed sediments of the Tethyan/High Himalayan realm; nearby sections expose Permian carbonates below the pass	275	metasedimentary cover of alternating quartzite-schist carbonaceous schist and marble	33.51	77.77	5328		-30 to 10				Pandey et al. 2020
	Khumbu Glacier	Quaternary glacier over High Himalayan Crystalline (gneiss–schist) with abundant leucogranite intrusions	1.5	Quartz–feldspar–muscovite–biotite gneisses and leucogranites; garnet/sillimanite locally.	27.96	86.93	4900		-5 to 35				Takeuchi et al, 2020; Kirkbride et al, 2023
	Zaskar Basin	Quaternary glacier within Zaskar; bedrock includes Greater Himalayan Sequence and Tethyan units	1.5	Quartz–feldspar–mica assemblages in gneiss/schist; calcite in carbonates	33.49	76.98	4500		-20 to 10				Thakkar et al, 2023
	Spiti Valley	Tethyan Himalaya marine succession from Precambrian–Cretaceous, notably rich Triassic–Jurassic–Cretaceous carbonates/shales.	100	Limestone/dolomite (calcite), shales with clays; ferruginous horizons locally	32.25	78.02	3800		-30 to 30				Mandal et al, 2021; Salwan et al, 2010; Singla et al., 2021
	Hunder, Shyok valley	Cenozoic Khardung volcanic suite (intermediate–felsic lavas) and intrusives of the Ladakh Batholith (Cretaceous–Paleogene); SSZ also exposes Shyok ophiolites/flysch	100	Volcanics rich in plagioclase–pyroxene ± quartz; batholith granodiorite (quartz–feldspars–biotite); ophiolite serpentized peridotite/gabbro assemblages.	34.59	77.42	3160		-20 to 35				Martin et al, 2025; Lone et al, 2023
	Nubra Valley	Quaternary aeolian dunes and glaciofluvial deposits over Ladakh Batholith / SSZ bedrock.	1.5	Dune/bar sands dominated by quartz with feldspar/mica; local carbonates in spring deposits	34.60	77.70	3048		-10 to 35				Chandra et al., 2023
	Pamir Plateau	Triassic–Cretaceous sedimentary sequences of South Pamir plus Early Cretaceous arc magmatism related to Neo-Tethys subduction	100	Arc volcanics (plagioclase–pyroxene–olivine ± amphibole) and sedimentary carbonates/siliciclastics (calcite, quartz–feldspar).	36.50	73.50	3000		-20 to 25				Liu et al., 2020;

Environment	Field site	Lithology (hosted rocks)	Geological age of host rock (Ma)	Associated Minerals	Lat	Long	Altitude (m)	pH	Temp (°C)	Salinity	Planetary Analog	Rationale	References
Saline ecosystems	Kyagar Tso Lake	Quaternary closed-basin lacustrine/playa deposits over Puga gneiss complex hosted in the Nidar Ophiolite	66	Evaporites incl. borax/ulexite (Na-borates), halite, gypsum/anhydrite, jarosite–alunite–copiapite–tamarugite, opaline silica, plus native sulfur reported in valley fill.	33.12	78.24	4700	9-10	-20 to 20	High	Icy Ocean Worlds	Subsurface brines under the ice, high salt concentrations, potential habitats in salt-rich oceans.	Pandey et al. 2020;
	Pangong Tso Lake	Quaternary–Holocene lacustrine system along India–Tibet suture, catchment on Ladakh Batholith and Karakoram	100	Carbonate muds, aragonite–calcite, with clastic quartz, feldspars; evaporite facies locally (halite, gypsum); hydrothermal carbonates along active faults.	33.72	78.90	4250	9	-35 to 20	Mid			Streule et al, 2009; Rathour et al, 2020; Sagwal et al 2023,
	Band-e-Amir Lakes	Lakes dammed by Holocene travertine; regionally underlain by Mesozoic (Jurassic–Cretaceous) carbonates in Bamyán	72	Travertine (CaCO ₃) forms the natural dams; surrounding bedrock is limestone/dolostone	34.84	67.23	3000	8-9	-10 to 25	Mid			Shroder et al., 2021;
	Tso Kar Lake	Holocene closed-basin saline lake deposits within Ladakh's Quaternary lacustrine systems hosted in Changthang	50	Evaporites including halite, gypsum, and borates (e.g., ulexite/borax) documented regionally in Ladakh basins.	33.30	78.00	4530	9-10	-20 to 20	High			Dey et al, 2023; Gopal et al, 2002; Ali et al, 2025
	Tso Moriri Lake	Late-glacial to Holocene lacustrine carbonates/siliciclastics; catchment on Trans-Himalayan/Tethyan units	50	Carbonate muds (calcite/aragonite) with clastic quartz–feldspar from catchment.	32.90	78.32	4522	8.5-9.5	-20 to 20	Mid			Sigoyer et al, 2004; Gopal et al, 2002
Tectono-magmatic terrains	Ladakh	Mesozoic (Jurassic to Lower Cretaceous) ophiolites	130	Olivine, pyroxene, serpentine (pseudomorphs of olivine and pyroxene), and chromite.	32.75	78.00	4100		-25 to 10		Lunar/Mars	Hyperarid conditions, high UV, serve as terrestrial analogs for surface environments on Mars and the Lunar regolith, where desiccation, intense radiation, and limited shielding define habitability constraints	Villalobos-Orchard et al. 2025; Bhat et al. 2022; Ullah et al. 2020
	Rangit Window	Tectonic window exposing Lesser Himalayan units incl. Gondwana (Permo-Carboniferous) within the Rangit duplex (bounded by Ramgarh Thrust/MBT)	300	Predominantly siliciclastic (quartz–feldspar–mica) with coal-bearing horizons in the Gondwana sequence	27.29	88.29	400		-5 to 27		Early Earth/Venus	Volcanic provinces, oceanic crustal rocks, and ductile crustal processes were widespread on the early Earth and Venus	Schopf et al., 2008

Table 4: Compilation of putative planetary analogue sites across the Indian Ocean classified by ecosystem type, geological context, mineralogy, and planetary analog relevance. Each site is described in terms of its host lithology, geologic age, and key geochemical parameters (pH and temperature), along with associated minerals. Collectively, these sites serve as valuable parallels to Early Earth conditions and icy ocean worlds. References for each field site correspond to published geological and geomicrobiological studies.

Environment	Field site	Location	Lithology (hosted rocks)	Geological age of host rock (Ma)	Associated Minerals	Lat.	Long.	Altitude (m)	pH	Temp (°C)	Planetary Analog	Rationale	References
Deep-sea trenches	Makran Trench	Arabian Plate subducting beneath Eurasian Plate	Active trench/accretionary wedge — sediment pile from Quaternary to Neogene turbidites and hemipelagic muds.	65	Detrital quartz–feldspar–mica from hinterland; calcite from biogenic carbonate ooze; clay minerals (illite/smectite).	10.00	66.00	-3500		2 to 4	Early Earth	Early Earth's deep ocean subduction environments where high pressure + nutrient fluxes influenced early biospheres.	Mouchot et al., 2010; Fisher et al., 2012; Chen et al., 2024
	Java (Sunda) Trench	Indo-Australian Plate subducting beneath Eurasian Plate	Active trench with Quaternary trench-fill turbidites over oceanic crust (~Cretaceous age)	45	Detrital quartz–feldspar–mica; volcanogenic glass; calcite/aragonite from biogenic sediment.	-10.32	109.97	-7290		1 to 2			Jamieson et al. 2022
	Diamantina Trench	Rift-related fracture zone	Cretaceous oceanic crust floored by basalt; trench-fill pelagic sediments (Neogene–Quaternary).	50	Basalt with plagioclase, pyroxene, olivine; pelagic carbonate ooze (calcite).	-35.00	104.00	-8047		1 to 2			Stewart and Jamieson, 2019; Zhang et al., 2024
Deep-sea vents	Wocan	Carlsberg Ridge	Active (modern) hydrothermal systems on slow- to intermediate-spreading ridges (Cretaceous–Jurassic oceanic crust).	100	Massive sulfide deposits — pyrite, chalcopyrite, sphalerite, barite, silica; basalts with plagioclase, clinopyroxene, olivine	6.37	60.52	-3105	n.d.	358	Icy Ocean Worlds	Hydrothermal activity at the ocean–rock interface inferred from plume chemistry; parallels Earth's vent-driven chemolithotrophy	Most et al., 2023
	Daxi	Carlsberg Ridge		100		6.80	60.30	-3450	n.d.	273			
	Tianxiu	Carlsberg Ridge		100		-3.65	63.75	-3500	n.d.	"high"			
	Onnuri	Central Indian Ridge		100		-11.42	66.42	-2170	n.d.	242			
	Kairei	Central Indian Ridge		100		-25.32	70.03	-2460	3.3	369			
	Solitaire	Central Indian Ridge		100		-19.55	65.83	-2606	4.4	307			
	Dodo	Central Indian Ridge		100		-18.33	65.47	-2745	3.3	356			
	Edmond	Central Indian Ridge		100		-23.88	69.60	-3320	3.1	382			
	Pelagia	South-East Indian Ridge		100		-26.15	71.43	-3690	3.2	368			
	Duanqiao	South-West Indian Ridge		100		-37.65	50.40	-1732	n.d.	277			
	Tiancheng	South-West Indian Ridge		100		-27.95	63.53	-2729	n.d.	"high"			
	Longqi	South-West Indian Ridge		100		-37.78	49.65	-2755	3.6	379			