

1 **An assessment of the quality of microanalysis of silicate glass using scanning**
2 **electron microscope-based energy dispersive spectrometry (SEM-EDS)**

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Abstract

The composition of volcanic glass records important clues into the origin and evolution of magmatic systems. However, the analysis of volcanic glass presents challenges when performed using electron-beam techniques, particularly due to Na mobility. While microanalysis of geological materials is usually performed using electron microprobe-based wavelength-dispersive spectrometry (EMP-WDS), we present here results of glass analysis using energy-dispersive spectrometry with a scanning electron microscope (SEM-EDS). We use three U.S. Geological Survey (USGS) whole-rock reference materials (i.e., RGM-1, STM-1, QLO-1) that were fused into glass using a double-fusion technique. Results for 12 sessions over two months using the SEM-EDS reveal excellent reproducibility, agreement with the expected values for RGM-1 and STM-1, and excellent counting statistics when using live acquisition times of 15 s; results for QLO-1 suggest that the fused material is somewhat inhomogeneous. For longer acquisition times (30, 60, 90 s), significant Na migration is observed; we conclude that minimal Na migration takes place with acquisition times of 15 s using our analytical instrumentation. We also compile data for RGM-1 obtained using the same instrument over the last decade, and we further demonstrate the reliability of SEM-EDS for microanalysis of silicate glasses. Comparison of our results with EMP-WDS results for a range of reference materials published by researchers at the USGS Alaska Volcano Observatory suggest that both EMP-WDS and SEM-EDS can yield adequate analysis of volcanic glass, but SEM-EDS can potentially lead to higher precision for major elements (1 wt.% concentration), while EMP-WDS typically leads to higher precision for minor elements. The higher precision of SEM-EDS results for major elements, combined with the more widespread distribution, lower usage cost, and higher ease of use of

such systems, makes analysis via SEM-EDS an attractive option for measuring major-element compositions of volcanic glass.

Paper highlight information

Analysis of volcanic glass is important in petrological studies, but it is also challenging from an analytical standpoint. While analysis is usually performed using the electron microprobe, cost of usage, the complexity of analytical protocols, and the relatively long acquisition times pose challenges to systematic studies of volcanic glass. With the advent of silicon drift detectors (SDD), analysis of Earth materials with energy dispersive spectrometry on a scanning electron microscope (SEM-EDS) presents an attractive alternative to the electron microprobe, providing faster, cheaper, and easier analysis. Yet, there is a need for systematic studies that assess the quality of analysis of volcanic glass using SEM-EDS. We present here a systematic study of three USGS reference materials that we fused into glass using a double-fusion technique. We present the results of a short-term study using all three materials, as well as selected data for reference material RGM-1 collected over the last decade in our laboratory. We also compare our SEM-EDS results with results from a recent study in which a large compilation of electron microprobe data from the USGS Alaska Volcano Observatory is presented. We show that SEM-EDS results have adequate precision, accuracy, and reproducibility, both in the short-term and long-term studies. In fact, we argue that results for major elements are potentially more precise than what is generally obtained using the electron microprobe.

56 **Keywords**

57 volcanic glass; electron-based microanalysis; energy-dispersive spectrometry; scanning electron
58 microscope; wavelength-dispersive spectrometry; electron microprobe

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Introduction

Volcanic glass can record important clues into the origin of volcanic rocks and the evolution of magmatic systems. Yet, measuring accurate volcanic glass compositions is often challenging using electron-beam microanalytical techniques due to beam damage during analysis, particularly for glass that is rich in Na and H₂O (e.g., Devine et al. 1995; Morgan and London 2005; Roman et al. 2006; among many others).

Major and minor-element microanalysis of geological materials is traditionally performed using electron microprobe-based wavelength dispersive spectrometry (EMP-WDS), which generally allows for quantification of elements in concentrations of hundreds of parts per million to tens of weight percent (e.g., Potts 1987; Reed 2005). However, due to the geometry and characteristics of WDS spectrometers, each individual EMP-WDS analysis typically takes tens to hundreds of seconds, leading to significant beam damage for beam-sensitive materials such as volcanic glass. Further, EMP-WDS employs relatively high beam currents (when compared to other electron-beam instruments), often as high as several μA , which further contributes to significant beam damage and Na migration when using EMP-WDS. Finally, the combination of relatively high cost of operation, more limited availability of EMP instrumentation, and relatively long analytical times (typically on the order of minutes) often makes it impractical to obtain large numbers of analyses per sample using EMP-WDS, which can be necessary to identify alteration and compositional heterogeneity in natural volcanic glass.

With the advent of solid-state Silicon drift detectors (SDD), the quality of scanning electron microscope-based energy dispersive spectrometry (SEM-EDS) analysis has improved

considerably. Ritchie et al. (2012) show that SEM-EDS analysis of glasses using SDD rivals the precision obtained with EMP-WDS for major ($>0.1\%$) and minor ($>0.01\%$) elements (see also Reed and Ware 1973; Newbury and Ritchie 2015). Further, individual SEM-EDS analyses are obtained much more quickly (~ 30 s) because all elements are measured simultaneously by a single detector. Finally, incident currents in SEM-EDS analysis are typically on the order of a few nA, up to 3 orders of magnitude lower than what is used in EMP-WDS analysis. Short analytical times, combined with lower incident currents, result in less beam-damage to the analyzed glass. Additionally, the much more widespread availability of SEM and the lower operating costs make it possible to obtain large numbers of glass SEM-EDS analyses quickly and affordably, allowing better characterization of compositional heterogeneity of volcanic glass.

One of the challenges to testing the quality of volcanic glass analyses is finding suitable reference materials that are homogenous and have a well characterized compositions that are similar to natural compositions. In this work, we analyze glasses that we synthesized in the laboratory from powdered U.S. Geological Survey (USGS) whole-rock reference materials – whose compositions are well characterized by independent methods – to assess the quality of SEM-EDS analysis of volcanic glass. We compare data obtained for three different glasses over 12 analytical sessions spanning a short period of time (~ 2 months). We also present and discuss analytical data for one of these glasses obtained over the last decade on the same instrument at Vanderbilt University.

100 **Materials and Methods**

101 *Materials*

102 USGS whole-rock powder reference materials RGM-1, STM-1, and QLO-1 were used in this
103 study (Flanagan 1976; Abbey 1983; Gladney and Roelandts 1988; Govindaraju 1994). RGM-1
104 corresponds to a rhyolite, STM-1 is a nepheline syenite, and QLO-1 is a quartz latite. Together,
105 these three reference materials span a wide range of compositions representative of silicic
106 magmas (Table 1).

107 Powders were synthesized into glass using a double-fusion technique, to ensure complete
108 homogenization and melting of the starting powders, under atmospheric pressure in a
109 Thermolyne 46100 benchtop muffle high-temperature furnace. Powders were transferred into
110 a Platinum crucible with a tight-fitting lid and placed in a preheated furnace at temperatures of
111 1000 °C for 1 hour, to limit volatilization of Na. Temperatures were then increased to 1600 °C
112 and held for two hours. Following the hold at 1600 °C, the crucibles were extracted from the
113 furnace and quenched in a bucket of ice water.

114 Glass beads retrieved after fusion were mounted in epoxy mounts, ground to expose bead
115 surfaces large enough for analysis (typically ~1 mm across) and polished to 0.5 µm grit. Mounts
116 were carbon-coated prior to analysis.

117 *Analytical method*

118 SEM-EDS analysis was performed using an Oxford AZtecEnergy Advanced Microanalysis System
119 equipped with an X-Max 50 mm² large-area Energy Dispersive SDD attached to a Tescan Vega3

Variable-Pressure SEM installed in the Department of Earth & Environmental Sciences at Vanderbilt University (EES-Vanderbilt). Analysis was performed using standard protocols employed in prior work focusing on volcanic glass at EES-Vanderbilt (Gualda et al. 2018, 2022; Pamukçu et al. 2020; Pitcher et al. 2021; Smithies et al. 2023; Chiaro et al. 2024; Harmon et al. 2024a, 2024b), which consist of:

- Accelerating voltage of 15 kV.
- Working distance of 15 mm, which is the optimized working distance for EDS analysis in the EES-Vanderbilt system.
- Beam intensity set to 17-20, yielding absorbed currents of ~2-5 nA.
- Analysis areas as rectangles of variable sizes, but with typical height and width of 10-50 μm (Figure 1).
- SDD process time of 4, to yield an optimum compromise between precision in the measurement of x-ray energies by the SDD detector and x-ray output count rates.
- Conditions chosen to yield dead times of ~50-60%, to maximize the x-ray output count rates in the EDS detector.
- Live acquisition times of 15 s.
- Pulse pileup and ZAF corrections applied during quantification using the Oxford software AZtec.
- Quantification using factory standard measurements, with frequent (daily to weekly) beam optimization performed on Cu tape; energy optimization – also using Cu tape – was performed following filament replacement.

141 • Si, Al, Ti, Fe, Mg, Mn, Ca, Na, and K quantified from the collected spectra; oxygen
142 calculated by stoichiometry.

143 • Final results presented in oxide weight percentages, normalized to 100% anhydrous.

144 Additionally, for STM-1 and QLO-1, we performed analyses using live times of 30, 60, and 90 s,
145 to test the effects of Na migration as a function of analytical time.

146 We collect populations of 15-30 individual analyses in succession for a given area of glass that is
147 initially assumed to be homogeneous in composition. We calculate the average and standard
148 deviation of each population to yield the best estimate of the glass composition and the
149 uncertainty associated with these values. Consideration of the results allows testing of the
150 initial assumption of homogeneity, and subpopulations can be defined when necessary. The
151 standard deviation we calculate for each population is a measure of the repeatability
152 measurement precision (i.e., the variability arising from multiple measurements on the same
153 sample in the same analytical session, expressed as 1 standard deviation)(JCGM 200 2012), and
154 it includes the effects of both analytical uncertainty and compositional variability of the fused
155 glass.

156 We identify and remove outlier values of individual oxides using the median and interquartile
157 range (IQR). Values outside the window of the median $\pm 1.5 \times \text{IQR}$ are automatically excluded
158 from the average and standard deviation calculations. In our studies of natural samples (see, for
159 instance Pamukçu et al. 2021), we completely exclude individual analyses with four or more
160 values identified as outliers (i.e., all oxides are excluded from the average and standard

deviation calculation). However, given the absence of crystals or vesicles in the fused glasses, no analyses were excluded in the present study.

For the sessions in which we tested longer analytical times (15-90 s), we calculate mean standard weighted deviation (MSWD) values (Wendt and Carl 1991) to characterize the extent of dispersion observed within each population relative to the analytical uncertainties. We use these values to assess the extent of Na migration in these populations.

Datasets

One of the datasets we present here (hereafter, “Short-term dataset”) consists of analyses of synthesized glasses STM-1, QLO-1, and RGM-1 obtained over 12 analytical sessions that took place over 2 months (October-November 2023), with 1-2 populations of data collected for each of the reference materials in each session (no RGM-1 analyses for the first two sessions). This allows us to evaluate the quality of the data generated using the SEM-EDS at EES-Vanderbilt for a range of compositions relevant for silicic and alkaline magmas.

We also present a compilation of a large quantity of data for RGM-1 collected over the last 10 years (hereafter, “Long-term RGM dataset”) using the SEM-EDS at EES-Vanderbilt (Pamukçu et al. 2020; Smithies et al. 2023; Chiaro et al. 2024; Harmon et al. 2024a, 2024b). This allows us to assess the long-term reproducibility of data generated using our instrument.

Data presented in the paper are available as a supplementary file available online.

Results

Short-term dataset

For live times of 15 s, the Short-term dataset includes 10 populations of RGM-1 glass (from analytical sessions 3-12), 13 populations of QLO-1 glass (two from session 11), and 12 populations of STM-1 (one from each session), as illustrated in Figure 2, Figure 3, and Table 1.

As explained above, we report average compositions calculated from 15-30 individual spot analyses collected during each session. The dataset also includes populations for live times of 30, 60, and 90 s for QLO-1 (session 2) and STM-1 (session 1), as shown in Figure 4.

Visual inspection of our results for live times of 15 s suggests excellent reproducibility for both STM-1 and RGM-1 (see Figure 2 and Figure 3), and also excellent agreement with the reference values for the respective whole-rock reference materials (shown as squares in Figure 2 and Figure 3). Variations in values of abundant oxides such as SiO₂ and CaO are very small, and they show excellent accuracy (as measured by the ratio of measured to expected values), particularly for RGM-1 (ranging from 99.6% to 100.0% for SiO₂ and 100.7%-106.0% for CaO) and STM-1 (100.0%-100.3% for SiO₂ and 101.3%-105.6% for CaO), with QLO-1 showing inferior accuracy (100.0%-100.9% for SiO₂ and 91.9%-98.0% for CaO). Variations in K₂O are larger, but accuracy is excellent: 98.9%-100.3% for RGM-1, 97.7%-99.6% for STM-1, and 94.9%-101.0% for QLO-1. Results for Na₂O show slightly inferior accuracy, but are still very good: 95.5%-99.8% for RGM-1, 97.0%-99.9% for STM-1, and 96.7%-105.1% for QLO-1. Uncertainties are significantly larger for oxides in lower abundances (e.g., TiO₂, MnO), but reproducibility and agreement with reference values are still observed. The QLO-1 results show much greater variability; given that

these measurements were performed in the same analytical sessions as the measurements for STM-1 and RGM-1, and the range of values measured are similar, we infer that the variability in QLO-1 results is due to heterogeneity in the fused glass.

Results obtained for STM-1 and QLO-1 using longer live times show much larger variation (Figure 4), increasingly so with increasing live time (i.e., variability is largest for 90 s, and it decreases progressively to 60, 30, and 15 s). For live times greater than 15 s, values for Na₂O are systematically lower than expected, while values for other abundant oxides (e.g., SiO₂, K₂O) are systematically higher than expected. MSWD values for Na₂O concentrations are below 1 (and below the expected maximum MSWD threshold; see Table 2) for acquisition times of 15 s, and they increase rapidly, particularly for STM-1, with increasing acquisition time (Table 2). This suggests that the dispersion observed for data obtained with acquisition times of 15 s is well explained by the analytical uncertainties; in contrast, for acquisition times of 30, 60, and 90 s, the populations are overdispersed (i.e., observed MSWD is larger than the upper MSWD threshold), suggesting that the spread of data within each population is larger than what can be explained by the analytical uncertainties. This suggests that data obtained with 15 s acquisition time avoids significant Na migration.

Relative standard errors (i.e., standard deviation divided by concentration) for each oxide correlate strongly with concentration (Figure 2, Figure 3, and Figure 7). For oxides with concentrations >10 wt.%, relative standard errors are generally <1%; for concentrations between 1 and 10 wt.%, relative standard errors are typically <10%; for concentrations <1 wt.%, errors are much larger, and they can surpass 100% relative for concentrations of 0.1 wt.% or less.

222 *Long-term RGM dataset*

223 Results for RGM-1 obtained at EES-Vanderbilt since 2014 show excellent reproducibility over
224 time (Figure 5, Figure 6), with uncertainties similar to those observed in the Short-term dataset
225 (see Figure 2, Figure 3). We do not identify any trends over time.

226 **Discussion**

227 *Quality of SEM-EDS analysis of natural silicate glasses*

228 Our results show several important features of SEM-EDS analyses of natural silicate glasses:

- 229 1. Glass damage by the electron beam – and resulting migration of Na away from the
230 beam – can be avoided by use of acquisition times of ≤ 15 s on our instrument under the
231 conditions used.
- 232 2. Measured compositions are in excellent agreement with expected values, given the
233 uncertainties estimated by repeated analysis of RGM-1 and STM-1 glasses within a
234 session; QLO-1 glass shows heterogeneity that can be explained by insufficient
235 homogenization during glass synthesis.
- 236 3. Precision obtained is excellent for major elements and adequate for minor elements,
237 with oxides in concentrations ≤ 0.05 wt.% having uncertainties large enough that
238 quantification is impractical.

239 These results are not surprising, in light of prior evidence suggesting that EDS analysis using
240 SDD is superior to EDS analysis using older technology (Ritchie et al. 2012). Further, our results
241 show that – for materials that are generally expected to be homogenous, such as volcanic glass

in pumice – repeated analysis of the material in a single session can yield appropriate estimates of uncertainty (specifically, estimates of intermediate precision) that include the effect of inhomogeneity of the analyzed material, in addition to counting statistics. This can not only lead to results that are more precise and accurate than those obtained using single spot analysis, but it also allows recognition of subtle differences and multiple compositional groups, as well as identification of materials that have suffered compositional changes due to alteration and weathering. Obtaining a larger number of analyses per sample results in much superior ability to assess compositional variability of volcanic glasses, which is more easily attainable with the shorter analytical times of SEM-EDS analyses (~15 s live time, ~30 s total) and lower operating costs compared with EMP-WDS analyses.

Comparison with EMP-WDS

In order to compare the quality of the SEM-EDS analyses obtained here with typical analysis by EMP-WDS, we compare the relationship between concentration and relative standard error we obtained for the Short-term dataset with the results of Loewen et al. (2023), obtained for a suite of glass reference materials used by the Alaska Volcano Observatory of the U.S. Geological Survey (AVO-USGS), as shown in Figure 7 (hereafter, “AVO dataset”). The advantage of this representation of the data is that it allows comparison of the precision achieved by the SEM-EDS and EMP-WDS, even though the specific glasses analyzed differ between the two laboratories.

For both datasets, we observe an approximately linear relationship in log-log space, suggesting a power-law relationship between concentration and relative standard error. Further, it is

striking that the SEM-EDS results show a steeper slope than the EMP-WDS results. This indicates that the precision obtained with SEM-EDS is potentially superior to that obtained using EMP-WDS for major elements in oxide concentrations greater than ~1 wt.%. For oxides in concentrations between ~0.5 and 1 wt.%, the two techniques yield results of similar precision. Finally, for minor elements with oxide concentrations lower than 0.5 wt.%, precision obtained with EMP-WDS is superior to that obtained with SEM-EDS.

These comparisons reveal that SEM-EDS is potentially preferable for analysis of major elements in natural silicate glasses. This can be explained by the combination of lower incident beam currents, use of a scanning beam, shorter acquisition times, better x-ray collection efficiency of the SDD detector compared to WDS detectors, and the lack of geometrical errors imposed by acquisition of data using a single detector in SEM-EDS analysis. It is useful to note that SEM-EDS analysis is performed using factory-supplied measurements of calibration materials (i.e., “standard-less”), which means that no calibration is necessary prior to analysis, rendering glass analysis by SEM-EDS much simpler and faster than analysis by EMP-WDS. Quality control of SEM-EDS analyses is performed by obtaining data for a secondary standard at the beginning of each analytical session, indicating that the instrument is performing adequately prior to analysis of any unknowns.

In contrast, EMP-WDS is preferable for analysis using electron-beam techniques of natural silicate glasses for minor elements, present in oxide abundances lower than ~0.5 wt.% and, particularly, for elements in oxide concentrations below 0.1 wt.% (Reed and Ware 1973). This is justified by the higher beam currents used (e.g., yielding higher counts), longer acquisition times, use of dedicated WDS detectors with much higher spectral resolution, and the small

effect of errors due to geometrical and beam-damage (i.e., Na-migration) on absolute oxide abundances measured (Reed 2005). Importantly, results obtained for major elements using EMP-WDS show very good precision, adequate for most applications.

One disadvantage of SEM-EDS analysis is that the total sum of oxides is not available, thus calculation of H₂O concentration by difference is typically not possible (at least in the case of “standard-less” analysis; see Geshi et al. 2021); it has been amply demonstrated that errors associated with such estimations are very large (unless specific correction methods are used; see Devine et al. 1995; Roman et al. 2006). Therefore, direct analysis of H₂O concentrations in silicate glasses by other methods is much preferred.

Implications

In this work, we show that both SEM-EDS and EMP-WDS can be successfully used for analysis of natural silicate glasses. For applications in which precision of major-element concentrations is paramount (e.g., glass geobarometry; see Gualda and Ghiorso 2014), SEM-EDS analysis is an excellent choice, possibly preferable to EMP-WDS. In contrast, for applications in which precision of minor elements is critical (e.g., Ti-in-quartz geochronometry; see Wark et al. 2007), use of EMP-WDS is preferred. For most applications, however, results obtained with both techniques are probably equally suitable. The much more widespread availability and lower usage costs of SEM-EDS systems stand as advantages when compared to EMP-WDS. Future work could explore the effectiveness of SEM-EDS analysis of minerals, in which spatial variability is expected, and effects of beam damage are much less severe or nonexistent.

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313 **Data availability**

314 All data used in the manuscript is presented within the manuscript, or in previous publications,
315 as cited.

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318 **References**

- 319 Abbey, S. (1983) Studies in “standard samples” of silicate rocks and minerals 1969-1982.
320 Canadian Geological Survey Paper, 83–15, 114 p.
- 321 Chiaro, G.R., Gualda, G.A.R., Miller, C., Giordano, G., and Morelli, C. (2024) Mush Architecture
322 and Processes in the Reservoirs of a Supereruption-Scale Magma System, Permian Ora
323 Ignimbrite (Northern Italy). *Journal of Petrology*, 65, egae016.
- 324 Devine, J.D., Gardner, J.E., Brack, H.P., Layne, G.D., and Rutherford, M.J. (1995) Comparison of
325 microanalytical methods for estimating H₂O contents of silicic volcanic glasses. *American*
326 *Mineralogist*, 80, 319–328.
- 327 Flanagan, F.J. (1976) Descriptions and analyses of eight new USGS rock standards. U.S.
328 Geological Survey Professional Paper, 840, 192 p.
- 329 Geshi, N., Yamasaki, T., Miyagi, I., and Conway, C.E. (2021) Magma chamber decompression
330 during explosive caldera-forming eruption of Aira caldera. *Communications Earth &*
331 *Environment*, 2, 200.
- 332 Gladney, E.S., and Roelandts, I. (1988) 1987 compilation of elemental concentration data
333 for USGS BHVO-1, MAG-1, QLO-1, RGM-1, SCo-1, SDC-1, SGR-1, and STM-1.
334 *Geostandards Newsletter*, 12, 253–362.
- 335 Govindaraju, K. (1994) 1994 compilation of working values and descriptions for 383
336 geostandards. *Geostandards Newsletter*, 18, 1–158.
- 337 Gualda, G.A.R., and Ghiorso, M.S. (2014) Phase-equilibrium geobarometers for silicic rocks
338 based on rhyolite-MELTS. Part 1: Principles, procedures, and evaluation of the method.
339 *Contributions to Mineralogy and Petrology*, 168, 1033.
- 340 Gualda, G.A.R., Grayley, D.M., Connor, M., Hollmann, B., Pamukçu, A.S., Begue, F., Ghiorso,
341 M.S., and Deering, C.D. (2018) Climbing the crustal ladder: Magma storage-depth
342 evolution during a volcanic flare-up. *Science Advances*, 4, eaap7567.
- 343 Gualda, G.A.R., Ghiorso, M.S., Hurst, A.A., Allen, M.C., and Bradshaw, R.W. (2022) A complex
344 patchwork of magma bodies that fed the Bishop Tuff supereruption (Long Valley
345 Caldera, CA, United States): Evidence from matrix glass major and trace-element
346 compositions. *Frontiers in Earth Science*, 10, 798387.
- 347 Harmon, L.J., Gualda, G.A.R., Grayley, D.M., Smithies, S.L., and Deering, C.D. (2024a) The
348 Whakamaru magmatic system (Taupō Volcanic Zone, New Zealand), part 1: Evidence
349 from tephra deposits for the eruption of multiple magma types through time. *Journal of*
350 *Volcanology and Geothermal Research*, 445, 107966.

- 351 Harmon, L.J., Smithies, S.L., Gualda, G.A.R., and Gravley, D.M. (2024b) The Whakamaru
352 magmatic system (Taupō Volcanic Zone, New Zealand), part 2: Evidence from ignimbrite
353 deposits for the pre-eruptive distribution of melt-dominated magma and magma mush.
354 *Journal of Volcanology and Geothermal Research*, 447, 108013.
- 355 JCGM 200 (2012) International Vocabulary of Metrology – Basic and general concepts and
356 associated terms (VIM), 3rd edition.
- 357 Loewen, M.W., Wallace, K.L., Lubbers, J., Ruth, D., Izbekov, P.E., Larsen, J.F., and Graham, N.
358 (2023) Glass electron microprobe analyses methods, precision and accuracy for tephra
359 studies in Alaska. Alaska Division of Geological & Geophysical Surveys Miscellaneous
360 Publication, 174, 20 p.
- 361 Morgan, G.B., and London, D. (2005) Effect of current density on the electron microprobe
362 analysis of alkali aluminosilicate glasses. *American Mineralogist*, 90, 1131–1138.
- 363 Newbury, D.E., and Ritchie, N.W.M. (2015) Performing elemental microanalysis with high
364 accuracy and high precision by scanning electron microscopy/silicon drift detector
365 energy-dispersive X-ray spectrometry (SEM/SDD-EDS). *Journal of Materials Science*, 50,
366 493–518.
- 367 Pamukçu, A.S., Wright, K.A., Gualda, G.A.R., and Gravley, D. (2020) Magma residence and
368 eruption at the Taupo Volcanic Center (Taupo Volcanic Zone, New Zealand): insights
369 from rhyolite-MELTS geobarometry, diffusion chronometry, and crystal textures.
370 *Contributions to Mineralogy and Petrology*, 175, 48.
- 371 Pamukçu, A.S., Gualda, G.A.R., and Gravley, D.M. (2021) Rhyolite-MELTS and the storage and
372 extraction of large-volume crystal-poor rhyolitic melts at the Taupō Volcanic Center: a
373 reply to Wilson et al. (2021). *Contributions to Mineralogy and Petrology*, 176, 82.
- 374 Pitcher, B.W., Gualda, G.A.R., and Hasegawa, T. (2021) Repetitive Duality of Rhyolite
375 Compositions, Timescales, and Storage and Extraction Conditions for Pleistocene
376 Caldera-forming Eruptions, Hokkaido, Japan. *Journal of Petrology*, 62, ega106.
- 377 Potts, P.J. (1987) *A handbook of silicate rock analysis*, 622 p. Chapman and Hall, New York.
- 378 Reed, S.J.B. (2005) *Electron Microprobe Analysis and Scanning Electron Microscopy in Geology*,
379 2nd ed. Cambridge University Press.
- 380 Reed, S.J.B., and Ware, N.G. (1973) Quantitative electron microprobe analysis using a lithium
381 drifted silicon detector. *X-Ray Spectrometry*, 2, 69–74.
- 382 Ritchie, N.W.M., Newbury, D.E., and Davis, J.M. (2012) EDS Measurements of X-Ray Intensity at
383 WDS Precision and Accuracy Using a Silicon Drift Detector. *Microscopy and*
384 *Microanalysis*, 18, 892–904.

- 385 Roman, D.C., Cashman, K.V., Gardner, C.A., Wallace, P.J., and Donovan, J.J. (2006) Storage and
386 interaction of compositionally heterogeneous magmas from the 1986 eruption of
387 Augustine Volcano, Alaska. *Bulletin of Volcanology*, 68, 240–254.
- 388 Smithies, S.L., Harmon, L.J., Allen, S.M., Gravley, D.M., and Gualda, G.A.R. (2023) Following
389 magma: The pathway of silicic magmas from extraction to storage during an ignimbrite
390 flare-up, Taupō Volcanic Zone, New Zealand. *Earth and Planetary Science Letters*, 607,
391 118053.
- 392 Wark, D.A., Hildreth, W., Spear, F.S., Cherniak, D.J., and Watson, E.B. (2007) Pre-eruption
393 recharge of the Bishop magma system. *Geology*, 35, 235–238.
- 394 Wendt, I., and Carl, C. (1991) The statistical distribution of the mean squared weighted
395 deviation. *Chemical Geology: Isotope Geoscience section*, 86, 275–285.
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398 **Figure captions**

399 *Figure 1. Example of backscattered electron (BSE) image showing analyzed areas for*
400 *quantitative analysis of the RGM-1 standard using the SEM-EDS at the Department of Earth &*
401 *Environmental Sciences, Vanderbilt University.*

402 *Figure 2. Binary diagrams showing results of analysis of reference materials STM-1, QLO-1, and*
403 *RGM-1 by SEM-EDS (small colored symbols), compared with expected values (large colored*
404 *squares). Average compositions of ~10-15 spot analyses collected during each analytical session*
405 *are shown as separate symbols, with uncertainties (i.e., the repeatability measurement*
406 *precision, expressed as 1 standard deviation) shown by black error bars. Results for STM-1 and*
407 *RGM-1 reproduce the expected values very well, with uncertainties increasing with decreasing*
408 *oxide abundance. For QLO-1, significant variability is observed, suggesting heterogeneity in the*
409 *fused glass.*

410 *Figure 3. Diagrams showing results of analysis of reference materials STM-1, QLO-1, and RGM-1*
411 *by SEM-EDS (small colored symbols) as a function of analytical session, compared with expected*
412 *values (large colored squares at session 0). Each small symbol represents the average*
413 *composition of ~10-15 spot analyses, with uncertainties (i.e., the repeatability measurement*
414 *precision, expressed as 1 standard deviation) indicated by black error bars. No trends over time*
415 *are seen, and results for STM-1 and RGM-1 show good agreement between our results and*
416 *expected values. For QLO-1, results for CaO, K₂O and FeO show much larger variability, with*
417 *some values falling outside what would be expected based on the reference values, suggesting*
418 *heterogeneity in the fused glass.*

419 *Figure 4. Binary diagrams showing the relationship between SiO_2 and Na_2O for individual*
420 *analyses using a wider range of analytical times for STM-1 and QLO-1. On the top diagrams,*
421 *individual analyses are shown as separate symbols, showing the range of results obtained for*
422 *each analytical time used. On the bottom diagram, the mean value for each analytical time is*
423 *shown, with error bars representing 1 standard deviation. Diagrams on the left show results for*
424 *STM-1, while diagrams on the right show results for QLO-1. For analytical times of 15 s (circles),*
425 *there is no significant trend in the diagram, showing that no significant Na loss takes place. For*
426 *times of 30 s (diamonds), 60 s (triangles), and 90 s (squares), there is a strong negative*
427 *correlation between SiO_2 and Na_2O , showing Na migration from the analyzed area, and*
428 *corresponding increase in SiO_2 . The longer the time, the larger the effect. For STM-1, in which*
429 *Na_2O concentrations are much higher than in QLO-1, the observed effect is also larger.*

430 *Figure 5. Binary diagrams showing results of analysis of reference material RGM-1 obtained*
431 *using the EES-Vanderbilt SEM-EDS since 2014 (labeled according to each study), compared with*
432 *expected values (gray squares with error bars). Results show excellent reproducibility, as well as*
433 *excellent agreement between our measurements and expected values.*

434 *Figure 6. Diagrams showing results of analysis of reference material RGM-1 by SEM-EDS (small*
435 *colored symbols) as a function of analytical session (ordered as a function of time, labeled*
436 *according to each study), compared with expected values (large gray squares, off scale to the*
437 *left of session 0). Agreement with expected values is excellent, and there are no evident trends*
438 *over time, showing the reliability of analysis of silicic glasses by SEM-EDS. Notice that, for some*
439 *of the data, MnO values were not quantified, and they are plotted at 0 concentration.*

440 *Figure 7. Log-log diagram showing the relationship between concentration and relative*
441 *standard error (two standard deviations – i.e., 2-sigma – divided by concentration). Colored*
442 *symbols show our results for reference materials STM-1, QLO-1, and RGM-1 (labeled Vanderbilt*
443 *dataset; see text for details), while gray dots represent data from Loewen et al. (2023) for a*
444 *large suite of reference materials (AVO dataset; see text for details). Uncertainties obtained by*
445 *both methods are similar, but SEM-EDS results show smaller uncertainties for concentrations >1*
446 *wt.%, while EMP-WDS results show smaller uncertainties for concentrations <0.5 wt.%.*
447

448 Tables

449 *Table 1. Expected values and associated uncertainties for reference materials STM-1, QLO-1,*
450 *and RGM-1.*

	STM-1		QLO-1		RGM-1	
	<i>Expected</i>	<i>Uncertainty</i>	<i>Expected</i>	<i>Uncertainty</i>	<i>Expected</i>	<i>Uncertainty</i>
SiO ₂	61.15	0.50	66.68	0.48	74.23	0.54
TiO ₂	0.14	0.01	0.63	0.03	0.27	0.02
Al ₂ O ₃	18.88	0.24	16.46	0.19	13.86	0.19
MgO	0.10	0.02	1.02	0.07	0.28	0.03
FeO	4.82	0.09	3.96	0.09	1.69	0.03
MnO	0.23	0.02	0.09	0.01	0.04	0.00
CaO	1.12	0.06	3.22	0.08	1.16	0.07
Na ₂ O	9.17	0.21	4.27	0.13	4.12	0.15
K ₂ O	4.39	0.07	3.66	0.12	4.35	0.10

452 *Table 2. MSWD values for populations of data for reference materials STM-1 and QLO-1*
453 *obtained with different acquisition times (15, 30, 60, and 90 s).*

Population	MSWD	n	MSWD upper threshold
STM-1 15 s	0.83	20	1.65
STM-1 30 s	3.65	20	1.65
STM-1 60 s	35.0	20	1.65
STM-1 90 s	175	20	1.65
QLO-1 15 s	0.44	20	1.65
QLO-1 15 s (2)	0.67	20	1.65
QLO-1 30 s	1.18	12	1.85
QLO-1 60 s	11.8	5	2.41
QLO-1 90 s	9.60	20	1.65

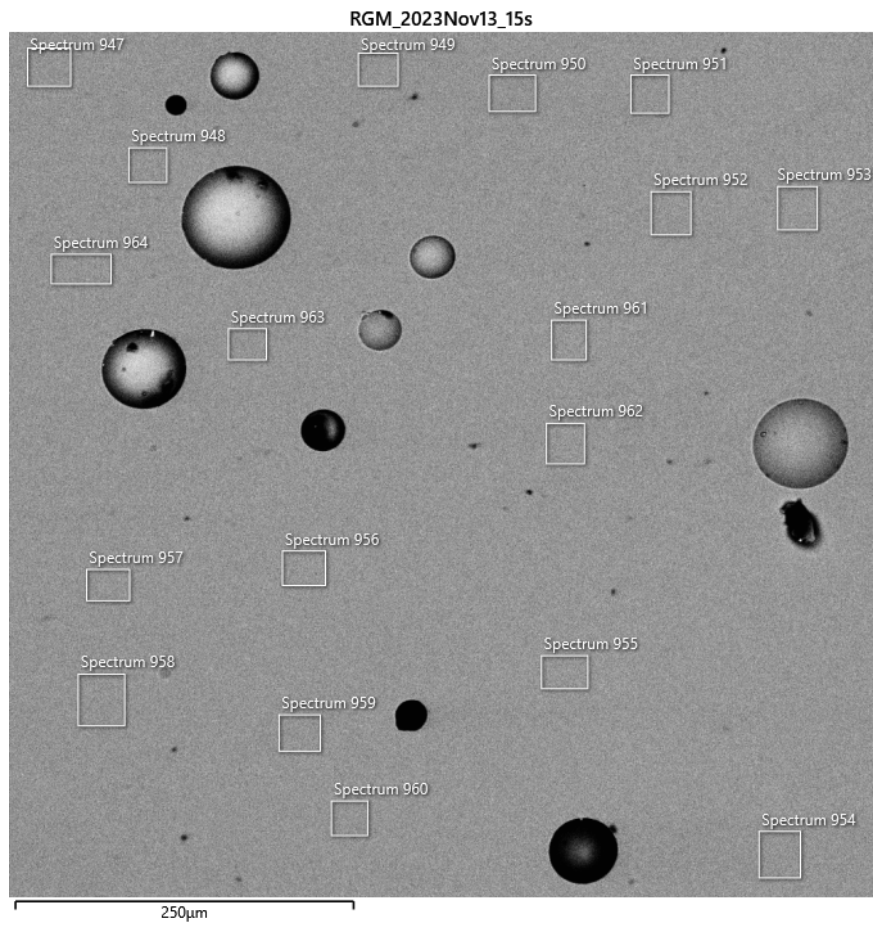


Figure 1

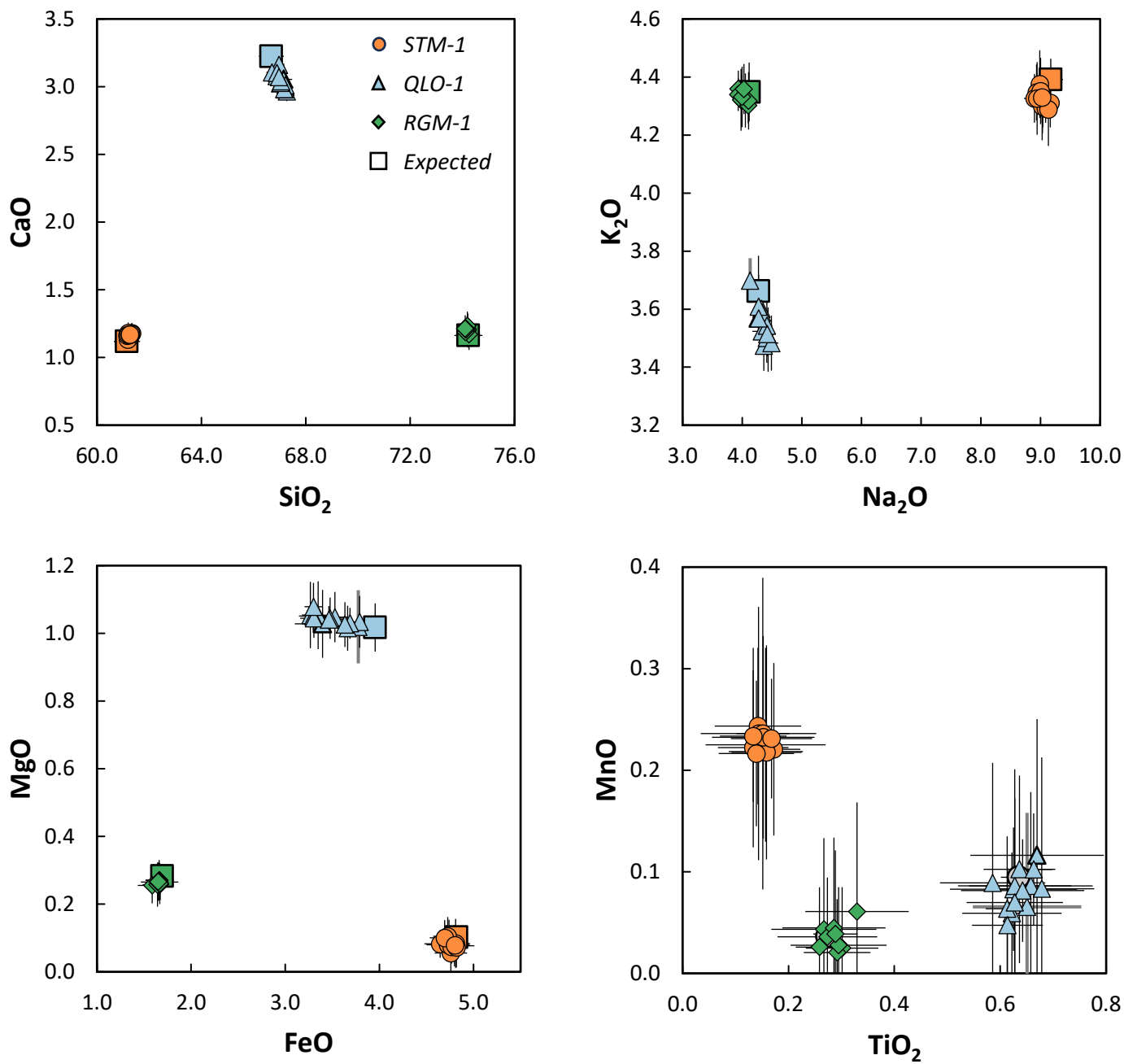


Figure 2

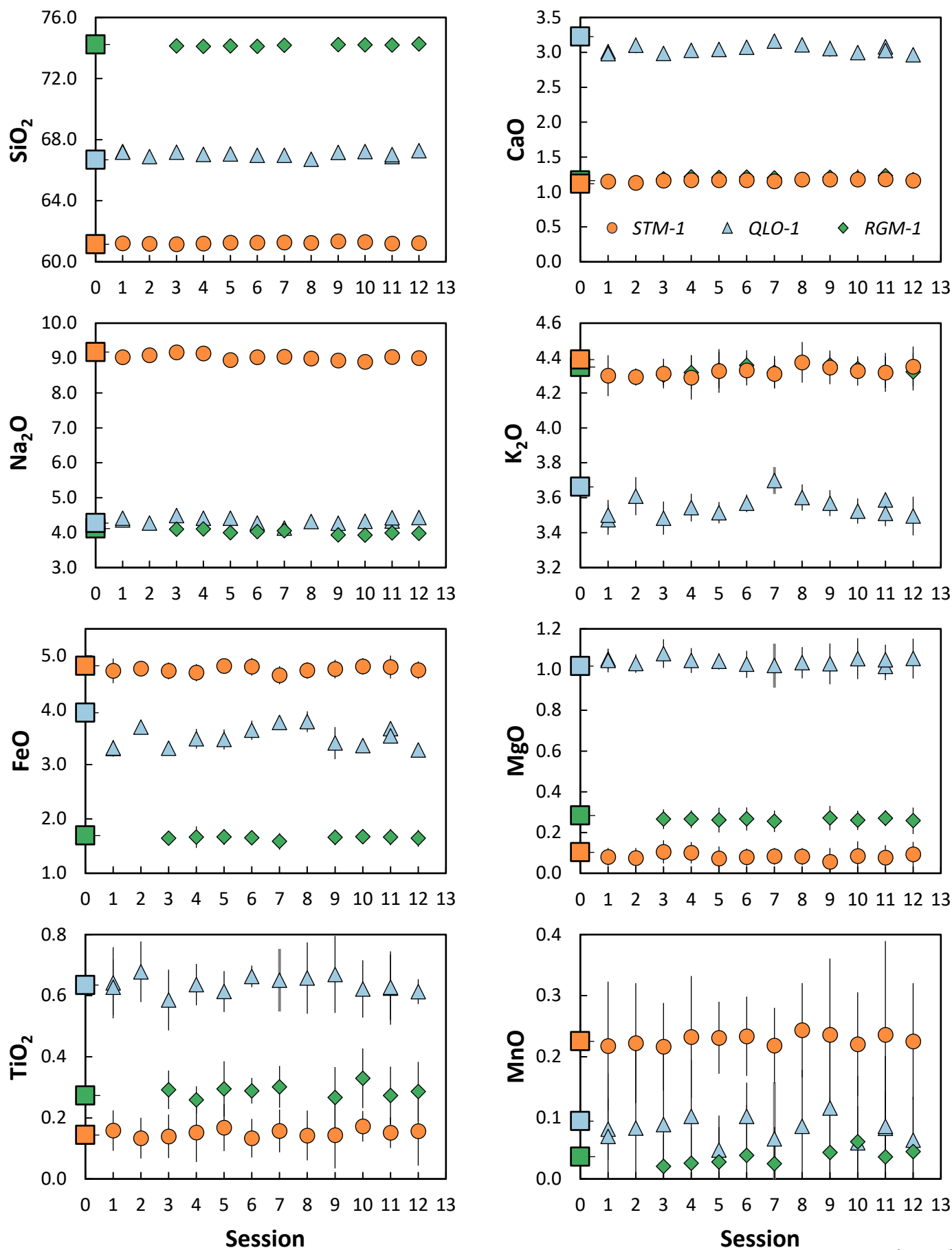


Figure 3

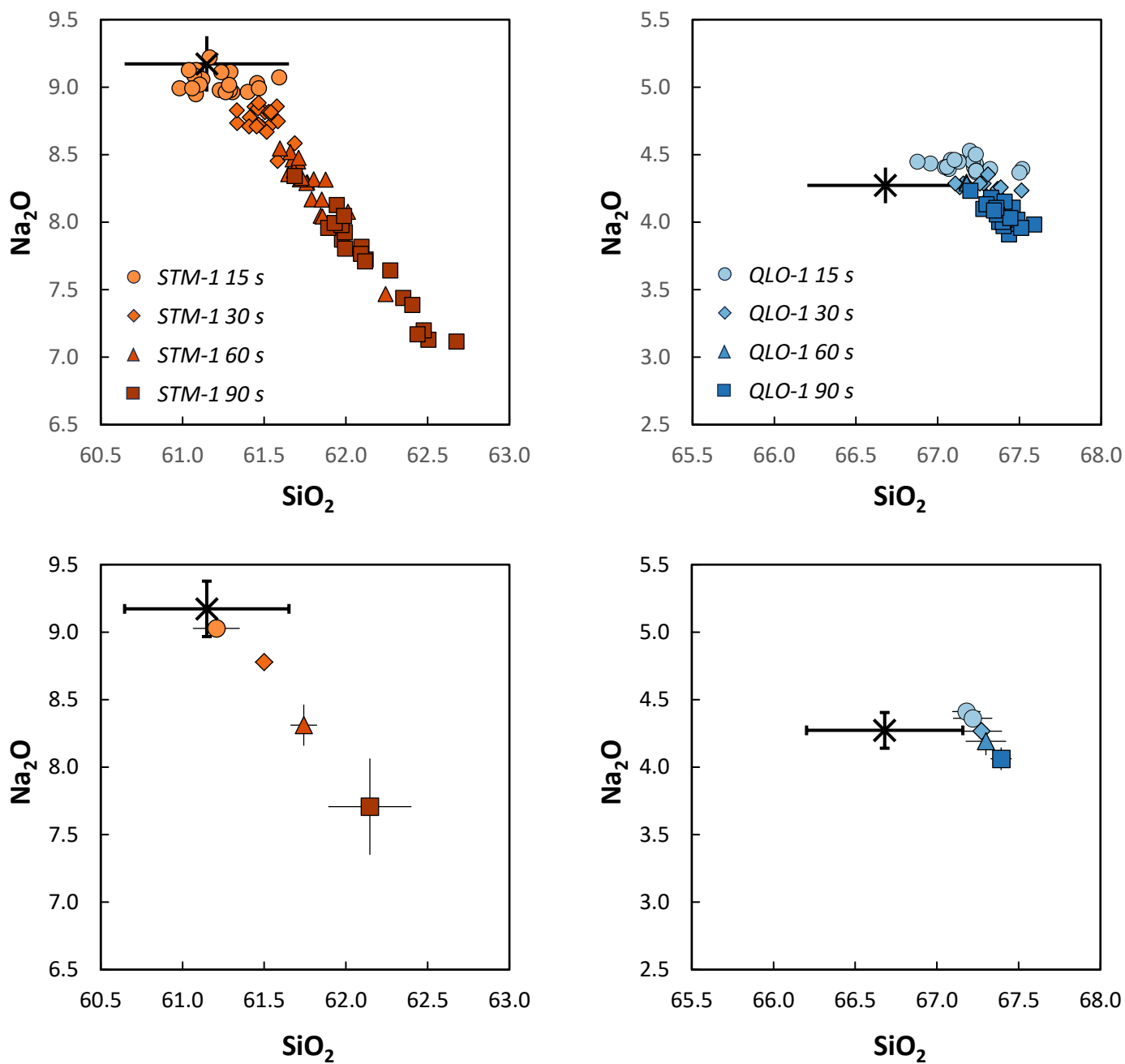


Figure 4

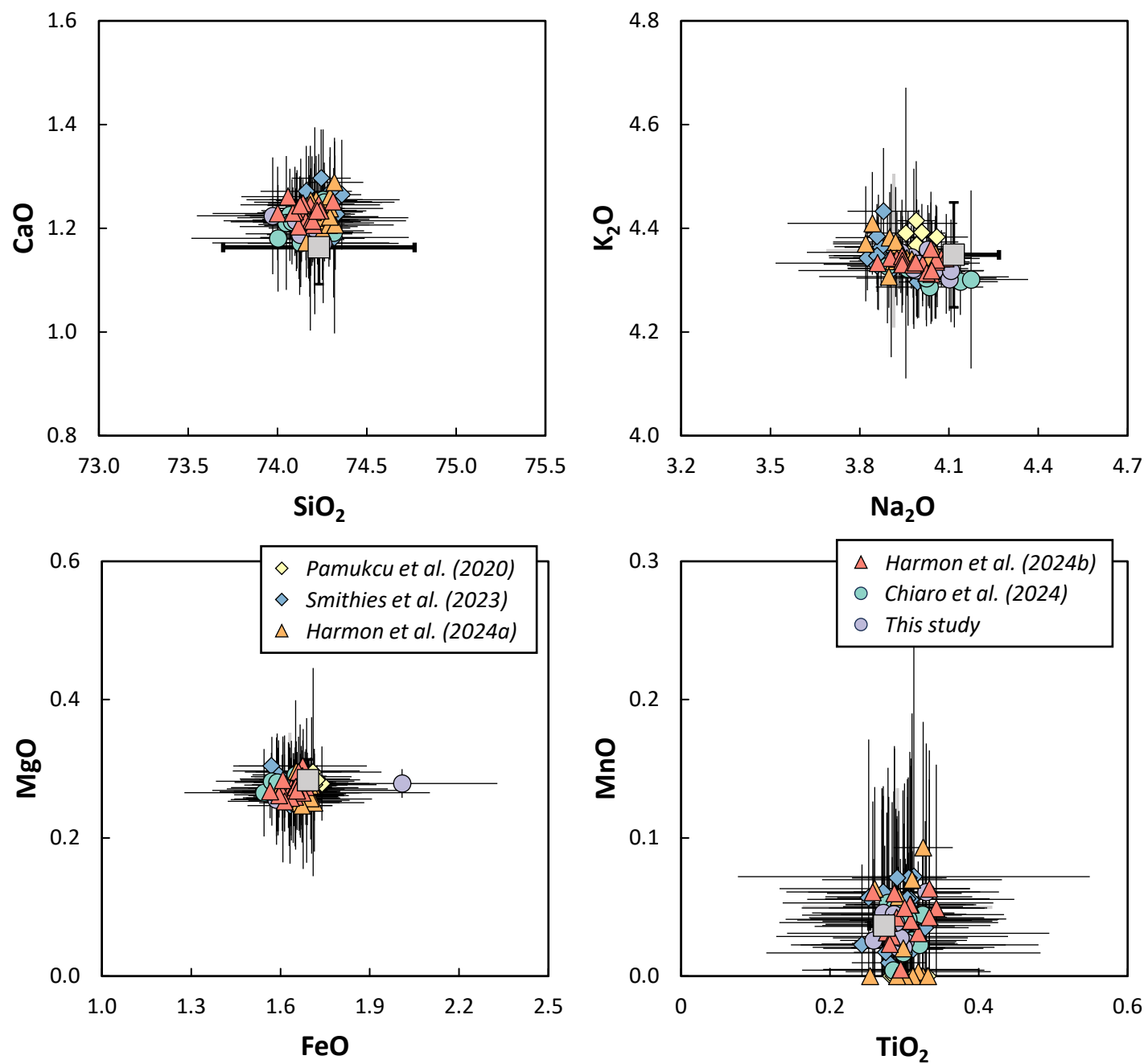


Figure 5

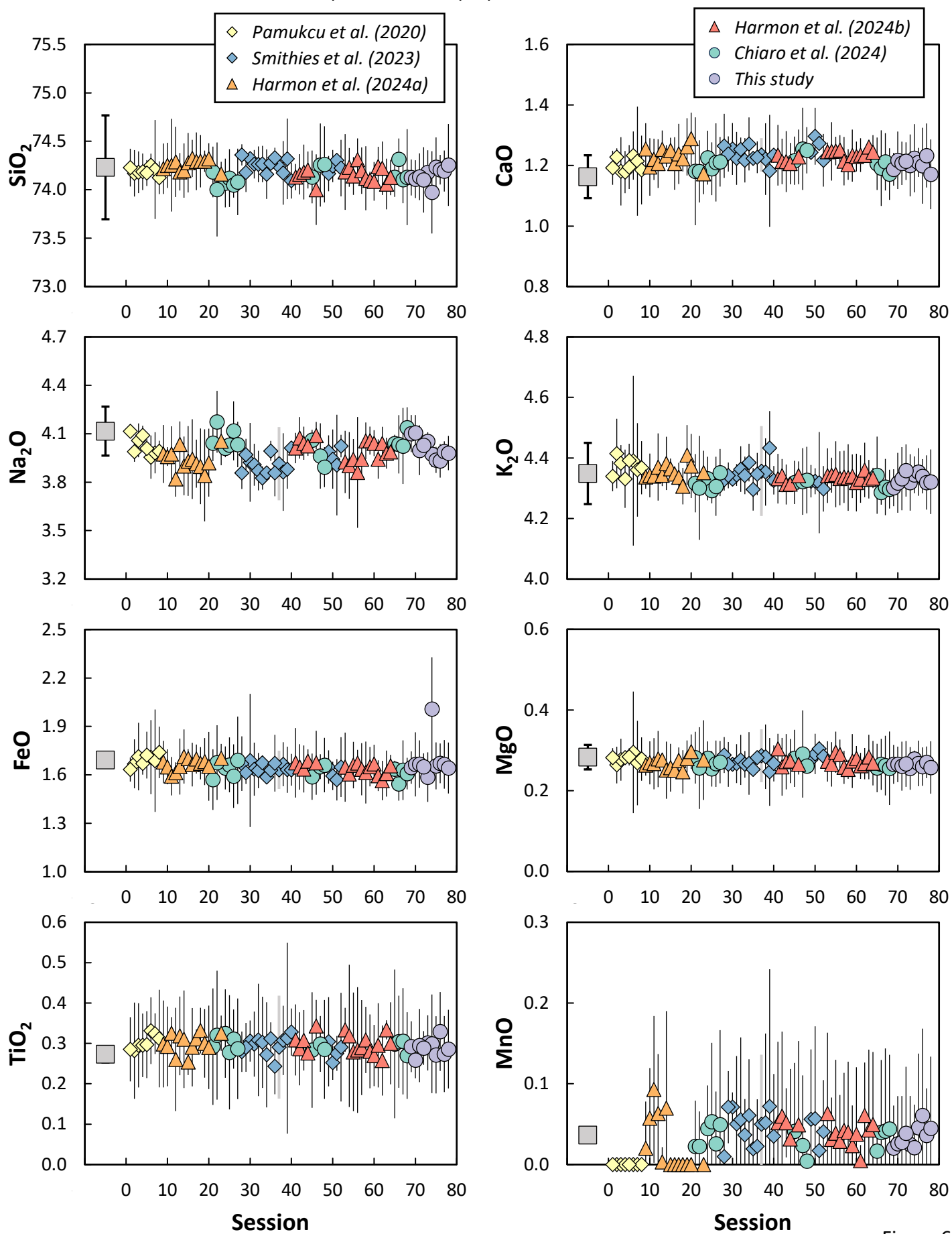


Figure 6

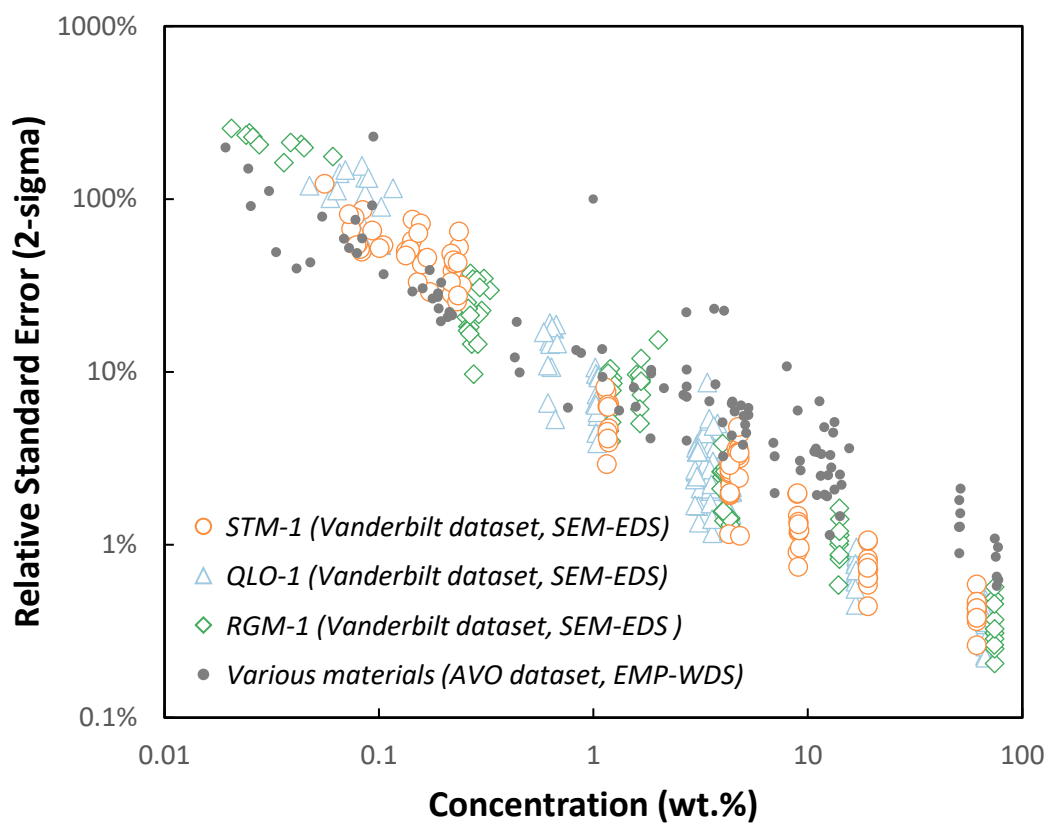


Figure 7