

# Spatial Predictor Selection for Next-Day Minimum Temperature Forecasting: An Automated Machine Learning Framework Applied Across European Climate Regimes

## Eric Duhamel

Correspondence: [edilia12380@gmail.com](mailto:edilia12380@gmail.com)

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## Abstract

9 Accurate prediction of daily minimum temperature (Tmin) is critical for frost protection, energy  
10 management, and public health preparedness. While numerical weather prediction models have  
11 improved substantially, their performance for Tmin forecasting remains limited by difficulties in  
12 representing fine-scale nocturnal processes. This study presents an automated framework for  
13 identifying optimal spatially-distributed predictors for next-day Tmin forecasting, applied to eight  
14 climatically diverse sites across Western Europe.  
15 Using 26 meteorological variables from ERA5 reanalysis data spanning 2004–2024, we systematically  
16 explored a search space of approximately 45,000 candidate predictors within a 540 km radius  
17 around each target station. An iterative optimization algorithm guided by mean absolute error  
18 (MAE) identified 90-predictor configurations for each site. Three regression models—linear  
19 regression, LightGBM, and XGBoost—were evaluated, with XGBoost consistently achieving optimal  
20 performance.  
21 Results demonstrate substantial skill across all sites, with MAE ranging from 0.81°C (Nice,  
22 Mediterranean) to 1.34°C (Brest, oceanic), representing 35–54% improvement over persistence and  
23 51–64% over climatological baselines. The analysis revealed both universal patterns—near-surface  
24 air temperature dominated predictive gain at all sites (37–66%)—and distinctive climate-specific  
25 signatures: Mediterranean stations exhibited strong persistence signals (30% contribution from  
26 previous-day Tmin), oceanic climates showed enhanced dewpoint importance (16%), and  
27 continental sites featured significant soil temperature contributions (14%).  
28 Predictor selections proved highly stable at the variable level (23–24 of 26 variables consistently  
29 selected across independent runs), while spatial autocorrelation caused greater variability in specific  
30 grid-point selections. Importantly, 80% of predictive gain originated from just 4 predictors and 90%  
31 from 12 predictors, suggesting that substantially reduced configurations could achieve comparable  
32 performance for operational applications.  
33 While the absolute MAE values reflect the idealized nature of reanalysis data and are not directly  
34 transferable to operational contexts, the methodology for predictor identification remains valid and  
35 applicable to numerical weather prediction outputs. This framework provides a systematic,  
36 reproducible approach to spatial predictor selection that can be adapted to other forecasting  
37 variables and domains.  
38 **Keywords:** *minimum temperature forecasting, spatial predictor selection, machine learning, ERA5*  
39 *reanalysis, gradient boosting, climate-specific signatures, Western Europe*

## 41 1. Introduction

### 42 1.1 Context and motivation

43 Forecasting daily minimum temperatures represents a major challenge in operational meteorology  
44 with significant implications across numerous sectors. Accurate predictions of the next day's  
45 minimum temperature are essential for protecting crops against frost, managing energy demand,  
46 winter road maintenance, and preparing public health services during cold spells (Wilks, 2011).  
47 Unlike maximum temperature forecasting, minimum temperature prediction is particularly sensitive  
48 to nocturnal radiative cooling, local topography, soil moisture conditions, and atmospheric boundary  
49 layer stability—factors that exhibit strong spatial variability and complex interactions (Stull, 1988).  
50 Numerical weather prediction (NWP) models, although constantly improving, often struggle to  
51 capture the fine-scale processes that control minimum temperatures, particularly in complex terrain  
52 or coastal areas. This limitation has motivated the development of statistical post-processing  
53 methods—often referred to as Model Output Statistics (MOS) (Glahn and Lowry, 1972)—capable of  
54 exploiting patterns in historical data to refine NWP model outputs or establish empirical  
55 relationships between atmospheric variables and surface temperatures.  
56 The spatial distribution of predictor variables represents a key but often under-explored dimension  
57 in temperature forecasting. While most statistical approaches focus on temporal structures or rely  
58 solely on local observations, atmospheric processes such as advection, regional cloud cover patterns,  
59 and maritime influences suggest that information from surrounding locations could significantly  
60 improve forecast quality. However, the optimal spatial configuration of predictors—which variables  
61 to use, at what distances, and in which directions from the target site—remains largely site-specific  
62 and poorly characterized.

### 63 1.2 Research objective and scope

64 **The primary objective of this study is to identify optimal sets of spatially distributed predictors for**  
65 **next-day minimum temperature forecasting at reference sites across Western Europe.** It is crucial  
66 to emphasize that this work focuses on **predictor selection methodology** rather than achieving the  
67 lowest possible prediction error in an operational context.

68 More specifically, we aim to:

- 69 1. Develop and validate an automated framework for systematic spatial exploration and  
70 predictor selection
- 71 2. Identify which meteorological variables, at what distances and in which directions from  
72 target sites, provide the most valuable information for  $T_{min}$  forecasting
- 73 3. Characterize how optimal predictor configurations vary across sites with contrasting climatic  
74 characteristics (maritime vs. continental, coastal vs. inland, northern vs. southern)
- 75 4. Quantify the added value of spatially distributed predictors compared to purely local  
76 approaches

77 This study uses a one-day temporal window to forecast the minimum temperature one day ahead,  
78 drawing on 26 meteorological variables derived from ERA5 reanalysis data (Hersbach et al., 2020)  
79 covering Western Europe, in addition to minimum temperature data from reference sites. The latter  
80 originate from NCEI-NOAA sources and correspond to actual observations. We apply this  
81 methodology to eight reference sites with diverse climatic characteristics (Peel et al., 2007), ranging  
82 from Mediterranean coastal sites to inland continental locations.

83 **It should be noted that the choice of minimum temperature as the target variable primarily serves**  
84 **to illustrate the methodological principle.** The developed approach is inherently generic and should  
85 apply without major difficulty to other meteorological variables (maximum temperature,  
86 precipitation, wind, humidity, etc.) or even to non-meteorological domains. For example, one can

87 envisage its application to river gauge level forecasting in hydrology, pollutant concentration  
88 forecasting in air quality, or any other environmental variable exhibiting spatio-temporal structure.

### 89 1.3 Methodological approach

90 Our approach combines systematic spatial exploration with machine learning regression techniques  
91 (Chen and Guestrin, 2016) (Ke et al., 2017) to automatically identify informative predictor sets. For  
92 each reference site, we explore the surrounding geographic space to extract candidate predictors  
93 from a set of 26 meteorological variables at different locations. These predictor sets are then  
94 evaluated using three complementary regression models—Linear Regression, LightGBM, and  
95 XGBoost—with mean absolute error (MAE) serving as the selection criterion.

96

97 **Mathematical formulation:** The regression problem can be expressed as finding the optimal  
98 function  $f$  such that:

99 
$$T_{min,t+1} = f(X_{1,t}, X_{2,t}, \dots, X_{n,t}) + \varepsilon_{t+1}$$
  
100 where  $X_i$  represents each predictor triplet (variable, latitude, longitude) at current day, and  $\varepsilon$  is the  
101 residual error. The objective is to minimize:  $MAE = (1/N) \sum |T_{obs} - T_{pred}|$   
102 where  $T_{obs}$  and  $T_{pred}$  denote the observed and predicted minimum temperatures, respectively, and  
103 the summation is taken over the  $N$  verification samples.

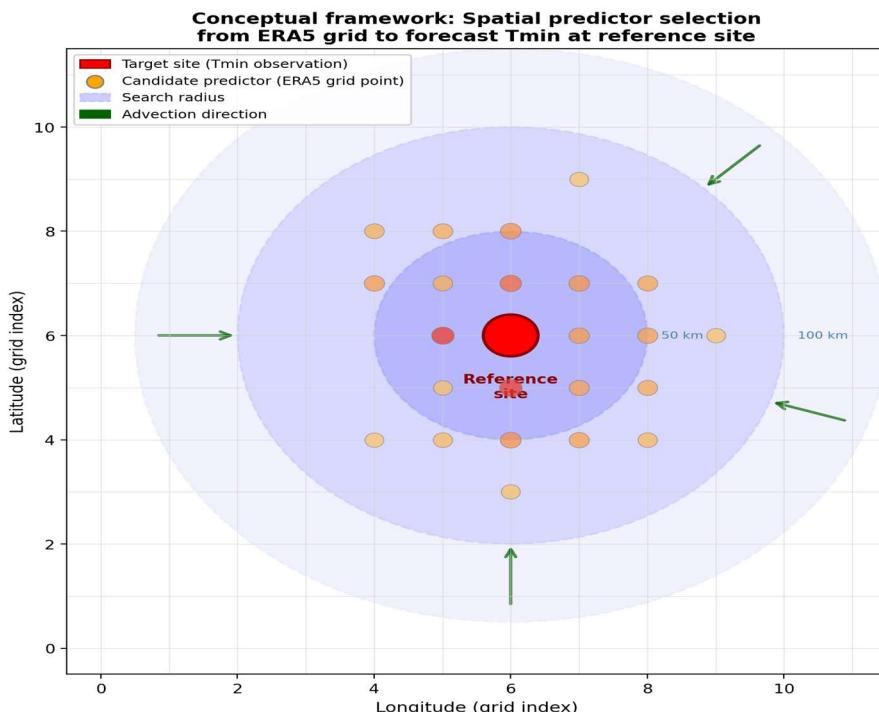
104

105 The use of multiple regression algorithms serves two purposes: (1) ensuring that identified  
106 predictors are robust regardless of the modelling approach, from linear to non-linear methods; and  
107 (2) providing insights into the nature of relationships between predictors and target temperatures.

108 To validate the added value of our spatial predictor selection approach, we implement rigorous  
109 comparisons with baselines including:

110 

- **Persistence models** (assuming tomorrow's temperature equals today's)
- **Climatological means** (historical average for each calendar day)



116 **Figure 1.** Conceptual framework illustrating the spatial predictor selection approach. The target site (red) is  
117 surrounded by candidate ERA5 grid points at varying distances (search radii). Point color intensity indicates

115 predictor relevance as determined by the selection algorithm. Arrows represent potential advection directions  
116 carrying information from distant locations.

## 117 1.4 Data limitations and scope of results

118 **A crucial caveat must be stated upfront regarding the interpretation of predictive performance**  
119 **metrics presented in this study.** All predictor data come from ERA5 reanalysis (Hersbach et al.,  
120 2020), while target observations (minimum temperatures) come from independent NCEI-NOAA  
121 station records. ERA5 represents a retrospective analysis that assimilates all available observations  
122 and applies sophisticated data assimilation techniques to produce an optimal, physically consistent  
123 reconstruction of past meteorological conditions.

124 **Consequently, the MAE values reported in this study are NOT representative of operational**  
125 **forecasting performance** that would be obtained using real-time NWP model outputs. ERA5 data  
126 benefit from:

- 127 • Complete spatial and temporal coverage with no missing data
- 128 • Assimilation of all available observations, including those not available in real-time
- 129 • Physical consistency optimized through retrospective analysis
- 130 • Absence of bias and drift inherent to forecast models

131 **Therefore—and this is the central thesis of our work—while absolute MAE values are not**  
132 **transferable to operational contexts, the methodology for identifying relevant predictors and their**  
133 **spatial configuration remains valid.** The physical relationships and spatial structures that emerge  
134 from our analysis reflect genuine atmospheric processes. The predictor sets identified through this  
135 methodology can be applied using operational forecast data, albeit with expected degradation in  
136 absolute performance.

137 This advantage of reanalysis-based predictors may, however, be partially counterbalanced in  
138 operational settings by the use of high-resolution short-range NWP forecasts, which can better  
139 capture local-scale processes despite their inherent forecast uncertainty. Additionally, a more  
140 judicious selection of features (variables) could also enhance the results.

## 141 1.5 Originality and contributions

142 This work makes several contributions to the field of statistical temperature forecasting:

- 143 1. **Systematic spatial exploration:** Unlike approaches that use predefined sets of stations or  
144 grid points, our methodology systematically explores geographic space to identify optimal  
145 predictor locations
- 146 2. **Multi-site comparative analysis:** By applying a consistent methodology to eight climatically  
147 diverse sites, we can characterize how optimal predictor configurations vary across local  
148 climate regimes
- 149 3. **Rigorous baseline validation:** Our comparison with persistence, climatology, and purely  
150 local models provides clear evidence of the added value of spatial predictor selection
- 151 4. **Physical interpretability:** The identified predictor sets can be interpreted in terms of known  
152 meteorological processes (advection, radiative effects, maritime influences)
- 153 5. **Open science and reproducibility:** All code, methodological details, and results will be made  
154 freely available via repositories such as GitHub and Zenodo, and Docker containers. The  
155 repository will be publicly released within 8 to 10 weeks of this preprint
- 156 6. **Approach genericity:** Although illustrated on minimum temperature, the methodology is  
157 applicable to other meteorological variables and other domains

## 158 2. Data

159 **2.1 Study area and reference sites**

160 This study focuses on Western Europe, covering a geographic domain bounded by latitudes 36.45°N  
161 to 63.16°N and longitudes 15.19°W to 18.44°E. This region encompasses a wide range of climatic  
162 conditions (Peel et al., 2007), from Mediterranean climates in the south to oceanic and semi-  
163 continental climates in the north and east, providing an ideal testbed for evaluating the spatial  
164 variability of optimal predictor configurations.



165  
166 **Figure 2.** Geographic distribution of the eight reference sites across Western Europe. The background shows  
167 the ERA5 grid domain used for predictor extraction.

168 Eight reference sites were selected to represent this climatic diversity (Table 1, Figure 2). The sites,  
169 located in the United Kingdom and France, include both coastal and inland locations, covering  
170 oceanic, semi-continental, and Mediterranean climate types.

171 **Table 1.** Reference sites used in this study.

Site	Lat. (°N)	Lon. (°E)	Country	Climate	Setting
Birmingham	52.42	-1.83	UK	Oceanic (Cfb)	Inland
Brest	48.44	-4.41	France	Oceanic (Cfb)	Coastal
Edinburgh	55.97	-3.21	UK	Oceanic (Cfb)	Coastal
Lyon	45.73	5.08	France	Semi-cont. (Cfb)	Inland
Nice	43.65	7.21	France	Medit. (Csa)	Coastal
Paris	48.72	2.38	France	Oceanic (Cfb)	Inland
Plymouth	50.35	-4.12	UK	Oceanic (Cfb)	Coastal
Strasbourg	48.55	7.64	France	Semi-cont. (Cfb)	Inland

172 **2.2 Reference data: observed daily minimum temperatures**

173 The target variable for this study—observed daily minimum temperature ( $T_{\min}$ )—was obtained from  
174 the NCEI-NOAA Daily Summaries database. Data were downloaded in CSV format for each of the  
175 eight reference sites.

176 The observation period spans 21 years, from 2004 to 2024 inclusive, providing 7,671 daily  
177 observations per site. This extended period captures a wide range of meteorological conditions,  
178 including extreme events, and ensures robust statistical analysis.

179 **It is essential to emphasize that these observed temperatures are entirely independent from the**  
180 **ERA5 predictor data described below.** This separation between target observations (station-based  
181 measurements) and predictor fields (reanalysis products) is fundamental to the validity of our  
182 predictor selection methodology.

183 **2.3 Predictor data: ERA5 reanalysis**

184 **What is a reanalysis?** A reanalysis is a systematic reprocessing of historical meteorological  
185 observations using a fixed, modern data assimilation system and numerical weather prediction  
186 model (Hersbach et al., 2020). Unlike operational forecasts that evolve over time as models are  
187 updated, reanalyses provide temporally consistent datasets by applying the same methodology  
188 throughout the entire period. ERA5, produced by the European Centre for Medium-Range Weather  
189 Forecasts (ECMWF), is the fifth generation of such reanalyses and currently represents the state of  
190 the art in global atmospheric reconstruction.

191 **Grid specifications:** ERA5 data are provided on a regular latitude-longitude grid with  $0.25^\circ \times 0.25^\circ$   
192 horizontal resolution (approximately 28 km at mid-latitudes). For our study domain [36.45°N–  
193 63.16°N, 15.19°W–18.44°E], this corresponds to approximately 107 latitude points  $\times$  135 longitude  
194 points, yielding roughly 14,000 grid cells.

195 **2.4 Predictor definition**

196 In this study, a predictor is formally defined as a triplet (variable, latitude, longitude) associated with  
197 a daily time series. For instance, "mean sea level pressure at grid point (48.25°N, 2.50°E)" constitutes  
198 a single predictor.

199 **Formal definition:** The total predictor space is defined as the Cartesian product:

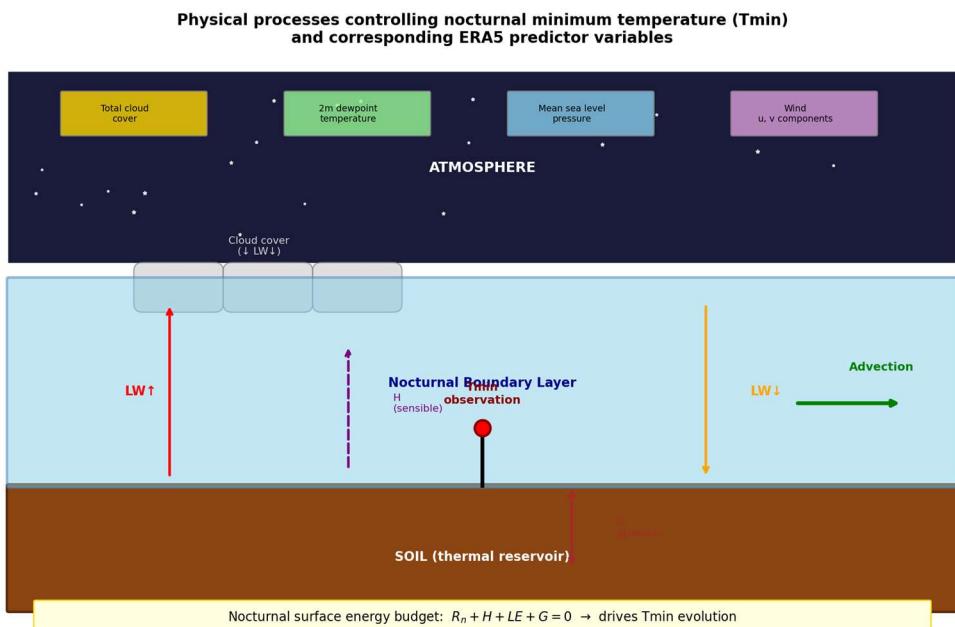
$$P = V \times \Lambda \times \Phi$$

200 where  $V = \{26 \text{ meteorological variables}\}$ ,  $\Lambda = \{107 \text{ latitude values}\}$ , and  $\Phi = \{135 \text{ longitude values}\}$ .

201 **Total candidate space:**  $|P| = 362,333$  potential predictors. This number is lower than the theoretical  
202 maximum ( $26 \times 107 \times 135 = 375,570$ ) because certain variable–location combinations are undefined  
203 (e.g., sea surface temperature over land, soil temperature over ocean).

205 **2.5 Predictor variables**

206 Daily minimum temperature ( $T_{\min}$ ) results from a complex interplay of physical processes, primarily  
207 nocturnal radiative cooling modulated by atmospheric and surface conditions (Oke, 1987). The  
208 predictor variables selected for this study were chosen to capture the main drivers of  $T_{\min}$  variability.  
209 They fall into six thematic categories:



210  
211 **Figure 3.** Schematic representation of physical processes controlling nocturnal minimum temperature ( $T_{min}$ ).  
212 The surface energy budget equation  $R_n + H + LE + G = 0$  governs temperature evolution, where  $R_n$  is net  
213 radiation,  $H$  is sensible heat flux,  $LE$  is latent heat flux, and  $G$  is ground heat flux. ERA5 predictor variables  
214 (colored boxes) are linked to these physical processes.

215 **Radiative balance:** Cloud cover (low, medium, high, total), downward thermal and solar radiation,  
216 net thermal radiation, and total column water vapour collectively determine the energy exchanges  
217 that drive nocturnal cooling (Stull, 1988). Low clouds and atmospheric moisture limit infrared losses,  
218 while clear skies promote strong radiative cooling.

219 **Advection and synoptic conditions:** Wind components ( $u, v$ ) and gusts characterize air mass origin  
220 and atmospheric mixing. Mean sea level pressure indicates synoptic regime—anticyclonic conditions  
221 favour calm, clear nights with low  $T_{min}$ , while cyclonic conditions bring clouds and milder nights.

222 **Boundary layer dynamics:** Boundary layer height (minimum and maximum) reflects atmospheric  
223 stability (Stull, 1988). A shallow nocturnal boundary layer promotes thermal inversions and radiative  
224 decoupling, leading to pronounced surface cooling.

225 **Surface and soil state:** Skin temperature and soil temperatures at two depths capture surface  
226 thermal inertia and heat exchanges with the atmosphere. Soil moisture modifies thermal  
227 conductivity and capacity. Snow depth, through its high albedo and insulating properties, strongly  
228 influences radiative cooling.

229 **Surface energy fluxes:** Latent and sensible heat fluxes, along with evaporation, complete the surface  
230 energy budget that governs temperature evolution from day to night.

231 **Thermal persistence:** The previous day's  $T_{min}$  and dewpoint temperature provide strong baseline  
232 predictors through temporal autocorrelation, with dewpoint additionally serving as a theoretical  
233 lower bound under saturation conditions.

234 **Table 2.** ERA5 predictor variables used for daily  $T_{min}$  forecasting.

Variable name	Description	Unit
10m_u_component_of_wind_daily_mean	Zonal wind at 10 m (daily mean)	m/s
10m_v_component_of_wind_daily_mean	Meridional wind at 10 m (daily mean)	m/s
10m_wind_gust_since_previous_post_processing_daily_max	Maximum wind gust at 10 m	m/s
2m_temperature_daily_minimum	Minimum air temperature at 2 m	K
2m_dewpoint_temperature_daily_minimum	Minimum dewpoint temperature at 2 m	K

boundary_layer_height_daily_minimum / maximum	Boundary layer height (daily min / max)	m
evaporation_daily_sum	Surface evaporation (daily accumulated)	m
low_cloud_cover_daily_mean	Low cloud cover (daily mean)	0-1
medium_cloud_cover_daily_mean	Medium cloud cover (daily mean)	0-1
high_cloud_cover_daily_mean	High cloud cover (daily mean)	0-1
total_cloud_cover_daily_mean	Total cloud cover (daily mean)	0-1
mean_sea_level_pressure_daily_mean	Mean sea level pressure (daily mean)	Pa
sea_surface_temperature_daily_mean	Sea surface temperature (daily mean)	K
skin_temperature_daily_minimum	Minimum surface skin temperature	K
soil_temperature_level_1_daily_minimum	Minimum soil temperature (layer 1)	K
soil_temperature_level_2_daily_minimum	Minimum soil temperature (layer 2)	K
surface_latent_heat_flux_daily_sum	Latent heat flux (daily accumulated)	J/m <sup>2</sup>
surface_sensible_heat_flux_daily_sum	Sensible heat flux (daily accumulated)	J/m <sup>2</sup>
surface_net_thermal_radiation_daily_sum	Net thermal radiation at surface (daily sum)	J/m <sup>2</sup>
surface_solar_radiation_downwards_daily_sum	Downward solar radiation at surface (daily sum)	J/m <sup>2</sup>
surface_thermal_radiation_downwards_daily_mean	Downward thermal radiation at surface (daily mean)	W/m <sup>2</sup>
total_column_water_vapour_daily_mean	Total column water vapour (daily mean)	kg/m <sup>2</sup>
total_precipitation_daily_sum	Total precipitation (daily accumulated)	m
volumetric_soil_water_layer_1_daily_mean	Volumetric soil water content (layer 1)	m <sup>3</sup> /m <sup>3</sup>
snow_depth_daily_mean	Snow depth (daily mean)	m

235

## 236 2.6 Potential redundancies

237 Several groups of variables exhibit expected collinearity. Cloud cover variables (total, low, medium,  
 238 high) are mechanically related, as are the various temperature fields (2 m air, skin, soil layers).  
 239 Radiative fluxes are partly determined by cloudiness and moisture content. These redundancies are  
 240 not necessarily detrimental: gradient boosting methods (e.g., XGBoost) and neural networks  
 241 typically handle correlated features without significant performance degradation (Chen and  
 242 Guestrin, 2016). However, for model interpretation—particularly when computing feature  
 243 importance or SHAP values—collinearity may dilute importance across redundant variables.

## 244 2.7 Data processing and storage

245 The predictor data (NetCDF format from ERA5) and reference observations (CSV format from NCEI-  
 246 NOAA) were integrated into a unified SQLite database to facilitate data management and querying.  
 247 This standardization offers several advantages:

- 248 • Uniform access to heterogeneous source formats through a single interface
- 249 • Efficient querying via SQL for data exploration and extraction
- 250 • Reproducibility through explicit data transformations
- 251 • Portability of the complete dataset as a single file

## 252 2.8 Objective and evaluation framework

253 A fundamental clarification is required regarding the role of performance metrics in this study. The  
 254 objective is *not* to minimize the mean absolute error (MAE) in absolute terms. Rather, the objective  
 255 is to **identify optimal sets of spatially distributed predictors**—that is, to determine which variables  
 256 at which geographic locations provide the most informative signal for  $T_{\min}$  forecasting.

257 **Evaluation metric:** The Mean Absolute Error (MAE) serves as a comparative criterion:

$$258 \quad MAE_{\text{config}} = (1/N) \sum_i |y_i - \hat{y}_i|$$

259 Configuration ranking:  $\text{config}_A > \text{config}_B \Leftrightarrow MAE_A < MAE_B$

260 In this context, the MAE serves as a **compass** rather than a destination. It provides a quantitative  
 261 criterion for comparing and ranking predictor configurations, guiding the search toward more

262 informative predictor sets. The absolute MAE value is less important than its *relative ranking* across  
263 configurations. A configuration yielding MAE = 1.5°C is preferred over one yielding MAE = 1.8°C,  
264 regardless of whether these values would be considered "good" or "poor" in an operational context.

265 This perspective has important implications for result interpretation. We prioritize the stability and  
266 interpretability of selected predictor sets over marginal performance improvements. A predictor  
267 configuration that consistently emerges across different model types and validation periods is more  
268 valuable than one that achieves slightly lower MAE but lacks robustness.

### 269 3. Methodology

270 This section describes the predictor selection methodology, including the definition of the search  
271 space, the selection algorithm, the optimization of the number of predictors, and the model  
272 configuration used for evaluation. The overall objective is not to achieve the best possible forecast  
273 accuracy *per se*, but rather to identify the most informative predictors for minimum temperature  
274 prediction. In this framework, the Mean Absolute Error (MAE) serves as a guiding criterion—a  
275 compass directing the search toward predictors that carry the strongest predictive signal (cf.  
276 Lakshmanan et al., 2015).

#### 277 3.1 Temporal Window Analysis

##### 278 Rationale

279 A critical methodological question in predictor selection for minimum temperature forecasting  
280 concerns the temporal depth of atmospheric information required (Box and Jenkins, 1976). While the  
281 atmospheric state on the current day provides direct predictive information for the minimum  
282 temperature on the following day, the inclusion of lagged variables from the previous two days could  
283 potentially capture synoptic-scale memory effects, particularly for continental stations where  
284 advective processes may exhibit multi-day persistence (Holmberg et al., 2024).

285 This experiment aimed to determine whether extending the temporal window beyond a single day  
286 improves predictive performance, or whether the atmospheric state at current day already encodes  
287 sufficient information, rendering lagged predictors redundant. The relevance of lagged predictors can  
288 be quantified through the temporal autocorrelation function. For a time series  $X$ , the autocorrelation  
289 at lag  $k$  is defined as:

$$290 \rho(k) = \text{Cov}(X_p, X_{p-k}) / \text{Var}(X)$$

291 where  $\rho(k)$  represents the correlation between observations separated by  $k$  time steps. High  
292 autocorrelation at lags  $k = 1$  or  $k = 2$  would suggest that including predictors from days D-1 or D-2  
293 might improve forecast skill.

##### 294 Experimental Design

295 Following the approach of lag selection in autoregressive modeling (Hyndman and Athanasopoulos,  
296 2021), three temporal window configurations were evaluated: (1) a single-day window (current day  
297 only); (2) a two-day window (incorporating a one-day lag); and (3) a three-day window  
298 (incorporating up to two-day lags). Each configuration was tested across four climatically contrasting  
299 stations: Birmingham (UK, oceanic), Brest (France, maritime), Nice (France, Mediterranean), and  
300 Strasbourg (France, continental).

301 For each station-window combination, the predictor selection algorithm was executed with five  
302 different random seeds to assess stability. Three machine learning models were applied: linear  
303 regression (as baseline), LightGBM, and XGBoost. Performance was measured using mean absolute  
304 error (MAE) on the test dataset (2023-2024), with results reported as mean  $\pm$  standard deviation  
305 across seeds.

306 Results were stored in JSON format with the following structure: for each window size (win\_size\_1,  
307 win\_size\_2, win\_size\_3), nested objects contain station-level results with model-specific MAE  
308 averages and standard deviations, plus mean elapsed computation time.

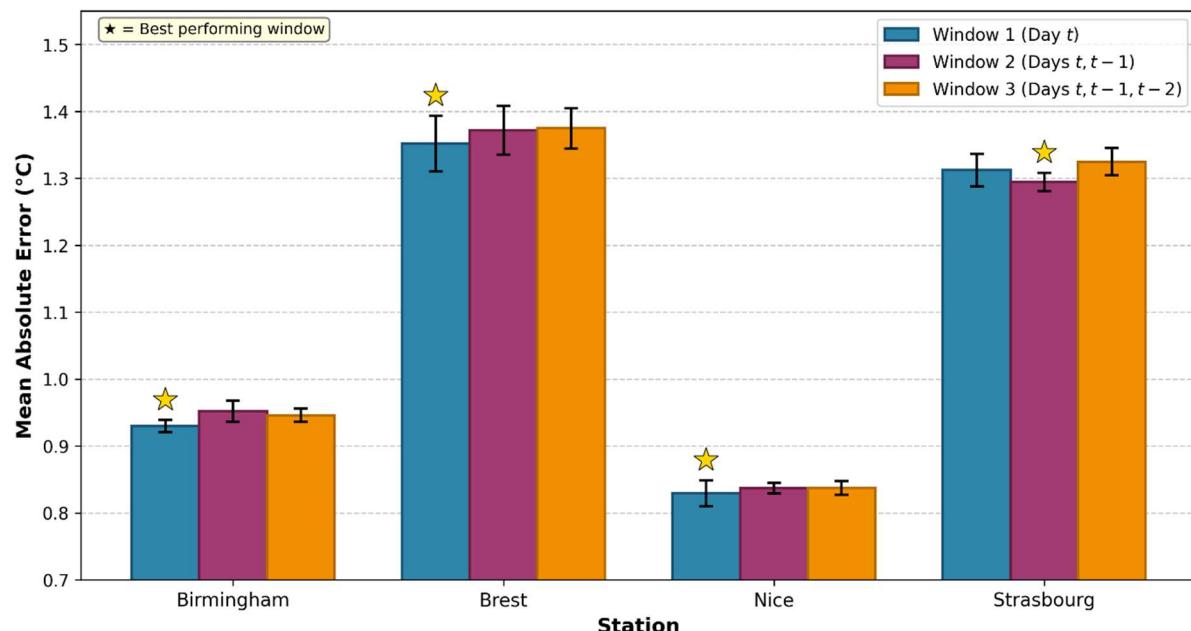
### 309 Results

310 Table 3 presents the MAE results for XGBoost, which consistently outperformed the other models  
311 across all configurations.

312 *Table 3. Mean absolute error (°C) by station and temporal window size (XGBoost model). Bold values  
313 indicate best performance. W1 = 1-day window, W2 = 2-day window.*

Station	Window 1	Window 2	Window 3	Best
Birmingham	<b><math>0.930 \pm 0.009</math></b>	$0.952 \pm 0.016$	$0.946 \pm 0.010$	W1
Brest	<b><math>1.352 \pm 0.041</math></b>	$1.372 \pm 0.037$	$1.375 \pm 0.030$	W1
Nice	<b><math>0.829 \pm 0.019</math></b>	$0.837 \pm 0.008$	$0.838 \pm 0.011$	W1
Strasbourg	$1.312 \pm 0.024$	<b><math>1.295 \pm 0.014</math></b>	$1.325 \pm 0.020$	W2
Avg. time	~40 min	~48 min	~55 min	

**Impact of Temporal Window Size on Minimum Temperature Prediction  
(XGBoost model, test period 2023-2024)**



314  
315 **Figure 4.** Comparison of MAE across temporal window sizes for each station. Error bars represent ±  
316 one standard deviation across five random seeds.

317

### 318 Analysis

319 The single-day temporal window achieved optimal or near-optimal performance for three of the four  
320 stations, consistent with the principle of parsimony in time series modeling (Burnham and Anderson,  
321 2002). Birmingham exhibited the clearest advantage for window 1, with a MAE reduction of 2.3%  
322 compared to window 2. Similar patterns emerged for Brest (~1.5%) and Nice (~1.0%). These results  
323 suggest that for oceanic, maritime, and Mediterranean climates, the atmospheric state on the current  
324 day contains sufficient predictive information for minimum temperature forecasting for the following  
325 day.

326 Strasbourg presented the sole exception, with the two-day window outperforming the single-day  
327 configuration by 1.3%. This finding aligns with the continental character of the station, where synoptic-  
328 scale processes—particularly cold air advections from northeastern Europe—may exhibit greater  
329 temporal persistence. The extended window likely captures the progressive establishment of high-

330 pressure systems and associated radiative cooling conditions that influence minimum temperatures  
331 over multiple days.

332 The three-day window consistently underperformed relative to shorter configurations, indicating that  
333 atmospheric information beyond the previous day introduces noise rather than useful signal.  
334 Standard deviations remained comparable across window sizes (typically 1-3% of MAE), suggesting  
335 that the stochastic variability inherent to the selection algorithm does not confound these  
336 comparisons.

337 Computation time increased approximately linearly with window size, with the three-day window  
338 requiring 35% more time than the single-day configuration. This overhead, combined with degraded  
339 predictive performance, reinforces the practical advantage of the parsimonious approach.

#### 340 Conclusion

341 The single-day temporal window was adopted as the default configuration for subsequent analyses.  
342 This choice reflects both empirical performance—optimal for 75% of stations tested—and the  
343 principle of parsimony advocated by Box and Jenkins (1976) for time series modeling. The atmospheric  
344 state at current day appears to encode the relevant synoptic and mesoscale information required for  
345 minimum temperature prediction, with lagged variables providing marginal or negative value.

346 This finding carries important implications for the physical interpretability of selected predictors:  
347 features identified under the single-day paradigm represent contemporaneous atmospheric drivers  
348 rather than lagged proxies, simplifying their meteorological interpretation and enhancing their  
349 potential transferability to operational forecasting contexts using numerical weather prediction  
350 outputs.

### 351 3.2 Predictor Search Space

352 The predictor selection algorithm explored a search space defined by the combination of ERA5  
353 variables (described in Section 2) and spatial locations within a fixed radius around each target  
354 station. ERA5 reanalysis (Hersbach et al., 2020) was deliberately employed to approximate an  
355 optimal atmospheric state, allowing the intrinsic predictive value of each variable to be assessed  
356 independently of operational forecast errors.

#### 357 Spatial domain

358 For each station, the candidate predictor pool comprised all ERA5 grid points within a radius  $R = 540$   
359 km. At the native ERA5 resolution of  $0.25^\circ$ , this radius encompasses a variable number of grid points  
360 depending on station latitude. Because ERA5 uses a regular latitude-longitude grid, the physical  
361 spacing between grid points in the zonal direction decreases with increasing latitude ( $\Delta x \approx 28 \text{ km} \times$   
362  $\cos(\phi)$ , where  $\phi$  is latitude). Consequently, higher-latitude stations capture more grid points within  
363 the same 540 km radius.

364 Across the eight stations, the total number of candidate predictors (grid points  $\times$  variables) ranged  
365 from approximately 41,000 (Nice,  $43.7^\circ\text{N}$ ) to 53,000 (Edinburgh,  $55.9^\circ\text{N}$ ), with a median of  
366 approximately 45,000. This latitudinal dependence reflects the convergence of meridians rather than  
367 any methodological inconsistency: the haversine distance metric ensures that all stations sample the  
368 same physical area ( $\pi \times 540^2 \approx 916,000 \text{ km}^2$ ), but the grid point density within that area varies with  
369 latitude.

370 The 540 km radius was chosen to capture synoptic-scale atmospheric patterns that influence local  
371 minimum temperatures (Holton and Hakim, 2013), while remaining computationally tractable.

#### 372 Temporal configuration

373 Predictors were extracted from ERA5 fields for the current day to forecast the overnight minimum  
374 temperature (i.e., the  $T_{min}$  for the following day). This configuration reflects an operational

375 forecasting scenario where atmospheric conditions observed during the day are used to predict the  
376 following night's minimum temperature.

377 In addition to the 26 ERA5 variables, the observed minimum temperature at the target station (from  
378 the NOAA GSOD dataset; Menne et al., 2012) **was included as a candidate predictor**. This allows the  
379 algorithm to exploit persistence—the tendency for consecutive days to have similar temperatures—  
380 where climatologically appropriate.

### 381 [3.3 Predictor Selection Algorithm](#)

382 Given the high dimensionality of the search space (ranging from ~41,000 to ~53,000 candidates  
383 depending on station latitude), exhaustive evaluation of all possible predictor combinations is  
384 computationally infeasible. The number of possible subsets of size  $k$  from  $n$  candidates is given by  
385 the binomial coefficient:

$$386 C(n, k) = n! / [k! \times (n - k)!]$$

387 Even for the smallest candidate pool ( $n \approx 41,000$  at Nice) and a modest subset size ( $k = 90$ ), this  
388 yields an astronomically large number of combinations ( $\sim 10^{200}$ ), rendering exhaustive search  
389 impossible. The latitude-induced variation in candidate pool size across stations (up to ~30%  
390 difference between Nice and Edinburgh) does not materially affect this conclusion, as all stations  
391 face comparably intractable combinatorial spaces. Instead, an iterative optimization algorithm was  
392 employed to identify near-optimal predictor subsets.

#### 393 [Objective function](#)

394 The algorithm sought to minimize the Mean Absolute Error (MAE) on a held-out validation dataset.  
395 The MAE is defined as:

$$396 MAE = (1/n) \times \sum_i |y_i - \hat{y}_i|$$

397 where  $y_i$  is the observed minimum temperature,  $\hat{y}_i$  is the predicted value, and  $n$  is the number of  
398 observations. MAE was preferred over Mean Squared Error (MSE) as it is less sensitive to outliers  
399 and provides a directly interpretable measure in °C (Willmott and Matsuura, 2005).

400 For each candidate predictor subset, a **linear regression** model was trained on the learning dataset  
401 and evaluated on the validation dataset. The MAE on the validation set served as the criterion  
402 guiding the search. Linear regression was chosen for its computational efficiency, enabling rapid  
403 evaluation of thousands of candidate subsets during the iterative search process.

404 At the end of the search process, the final predictor set was evaluated using gradient boosting  
405 models (LightGBM and XGBoost), which achieved lower MAE than linear regression due to their  
406 ability to capture non-linear relationships. The systematic use of gradient boosting during the search  
407 would have substantially increased computational time, hence the two-stage approach: linear  
408 regression for search guidance, gradient boosting for final evaluation.

#### 409 [Search strategy](#)

410 The feature search algorithm iteratively explored the predictor space, evaluating candidate subsets  
411 and progressively refining the predictor set toward lower validation MAE. The search was guided by  
412 heuristics that balance exploration of new predictor combinations with exploitation of promising  
413 solutions identified in previous iterations. This approach belongs to the family of metaheuristic  
414 optimization methods, which have proven effective for high-dimensional subset optimization  
415 problems (Xue et al., 2016).

#### 416 [Stopping criterion](#)

417 The algorithm employed an early stopping criterion based on lack of improvement. If no new best  
418 solution (lower validation MAE) was found within  $P = 60$  consecutive iterations, the search  
419 terminated and returned the predictor subset corresponding to the lowest validation MAE

420 encountered during the entire run. This patience-based stopping criterion (Prechelt, 1998) balances  
421 thorough exploration against computational cost.

422 **Stochastic variability**

423 Due to the stochastic nature of the search algorithm, different runs with different random seeds  
424 generally converge to different predictor subsets. To assess the robustness of the results, five  
425 independent runs were performed for each station, each initialized with a different random seed.  
426 This allows analysis of both the consistency of predictor selections across runs and the variability in  
427 prediction performance.

428 **3.4 Predictor Count Optimization**

429 **3.4.1 Rationale**

430 The number of predictors retained by the search algorithm represents a critical trade-off between  
431 model expressiveness and parsimony. Too few predictors may fail to capture the full complexity of  
432 atmospheric drivers influencing minimum temperatures, while excessive predictors risk introducing  
433 redundant information, increasing computational cost, and potentially degrading generalization  
434 through overfitting.

435 **3.4.2 Experimental Design**

436 Four predictor counts were evaluated: 50, 70, 90, and 110. Based on the results of Section 3.1, all  
437 experiments used a single-day temporal window (current day only). Tests were conducted across the  
438 same four climatically contrasting stations: Birmingham, Brest, Nice, and Strasbourg. For each  
439 configuration, the predictor selection algorithm was executed with five different random seeds, and  
440 performance was evaluated using XGBoost on the test dataset (2023–2024). Mean computation time  
441 was recorded to assess the computational cost of each configuration.

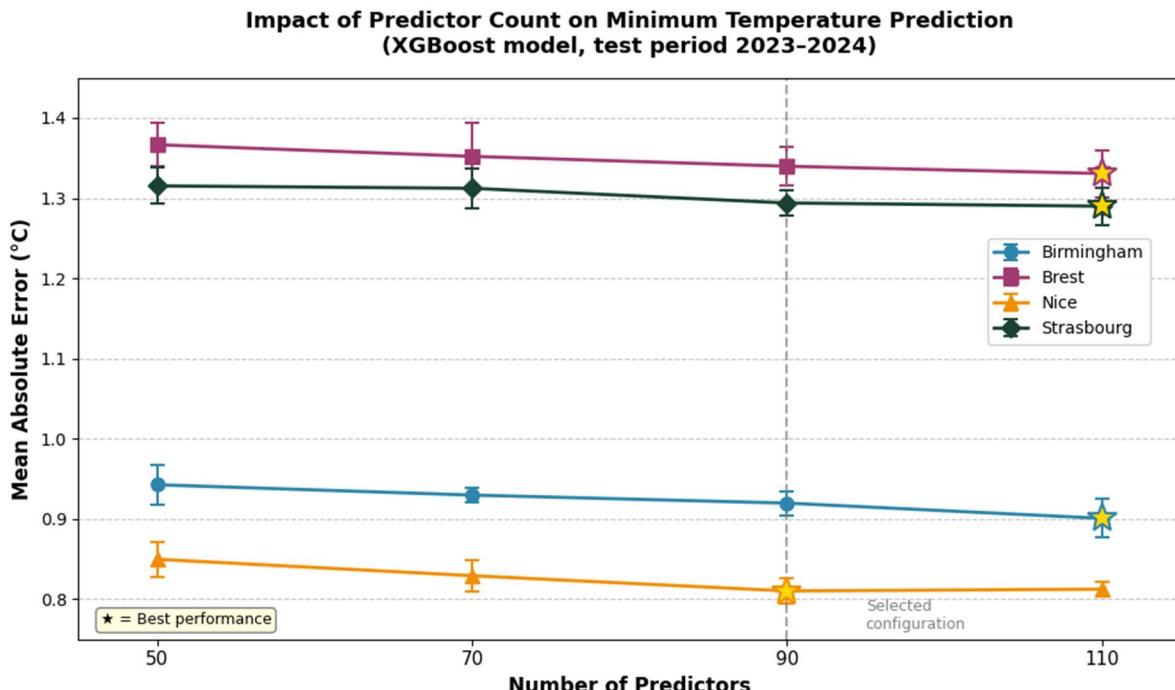
442 **3.4.3 Results**

443 Table 4 presents the MAE results for each station and predictor count, along with the relative  
444 improvement between successive configurations.

Station	50	70	90	110	Best
Birmingham	0.943	0.930	0.920	<b>0.901</b>	110
Brest	1.367	1.352	1.340	<b>1.331</b>	110
Nice	0.850	0.829	<b>0.810</b>	0.813	90
Strasbourg	1.315	1.312	1.294	<b>1.290</b>	110
Avg. time	~23 min	~40 min	~60 min	~78 min	

445 *Table 4. Mean absolute error (°C) by station and number of predictors (XGBoost model). Bold values indicate  
446 best performance for each station.*

447



448

449 **Figure 5.** Sensitivity of prediction accuracy to predictor subset size (XGBoost model, test period 2023–  
450 2024). The vertical dashed line indicates the selected configuration (90 predictors); stars denote best  
451 performance per station.

452

#### 453 3.4.4 Analysis

454 The results reveal a consistent pattern of diminishing returns as predictor count increases. All stations  
455 exhibited substantial improvement when moving from 50 to 70 predictors, with relative MAE  
456 reductions ranging from 0.2% (Strasbourg) to 2.5% (Nice). The transition from 70 to 90 predictors  
457 yielded further gains across all stations, ranging from 0.9% (Brest) to 2.3% (Nice).

458 Beyond 90 predictors, the behavior diverged across stations. Birmingham continued to benefit from  
459 additional predictors, with a 2.1% improvement at 110 predictors—the largest marginal gain observed  
460 in the 90-to-110 transition. This suggests that the oceanic climate of central England, influenced by  
461 multiple synoptic-scale drivers, requires a larger predictor set to capture its full complexity. Brest and  
462 Strasbourg showed marginal improvements (0.7% and 0.3%, respectively), while Nice exhibited a  
463 slight performance degradation (+0.3%), indicating potential overfitting when the predictor count  
464 exceeds the intrinsic dimensionality of the prediction problem.

465 The relationship between predictor count and computational cost proved approximately linear, with  
466 execution time increasing from ~23 minutes at 50 predictors to ~78 minutes at 110 predictors. This  
467 represents a 3.4-fold increase in computation time for a modest improvement in predictive  
468 performance.

469 Standard deviations remained stable across configurations (typically 1.5–2.5% of MAE), indicating that  
470 the search algorithm achieved consistent convergence regardless of the predictor count constraint.

#### 471 3.4.5 Conclusion

472 Based on these results, 90 predictors was selected as the default configuration for subsequent  
473 analyses. This value represents the point of diminishing returns for three of the four stations tested:  
474 beyond this threshold, performance gains become marginal (Brest, Strasbourg) or negative (Nice),  
475 suggesting that additional predictors introduce redundancy or noise rather than useful information.

476 The choice of 90 predictors also offers a practical compromise between predictive accuracy and  
477 computational cost. Compared to the 110-predictor configuration, it reduces execution time by  
478 approximately 20% while sacrificing less than 1% in MAE for most stations.

479 Although Birmingham exhibited continued improvement at 110 predictors, a uniform configuration  
480 was retained for methodological consistency. This facilitates cross-station comparison of selected  
481 predictors and avoids introducing station-specific tuning that could complicate interpretation. The  
482 modest performance difference ( $0.019^{\circ}\text{C}$ ) was deemed insufficient to justify a heterogeneous  
483 approach.

### 484 **3.5 Model Configuration**

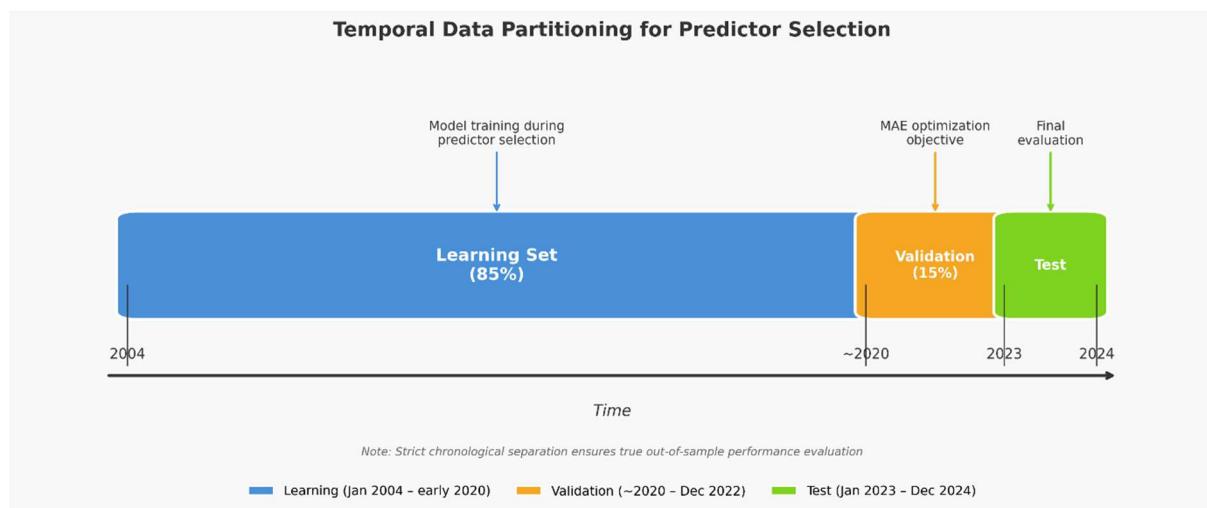
#### 485 **Data partitioning**

486 The available data (2004–2024) was partitioned chronologically into three non-overlapping periods,  
487 following best practices for time series forecasting (Bergmeir and Benítez, 2012):

488 **Learning set (85%):** January 2004 to approximately early 2020. Used for training the regression  
489 models during predictor selection.

490 **Validation set (15%):** Approximately 2020 to December 2022. Used to evaluate candidate predictor  
491 subsets and guide the selection algorithm. The MAE on this set served as the optimization objective.

492 **Test set:** January 2023 to December 2024. Held out entirely during predictor selection and used only  
493 for final performance evaluation. This strict temporal separation ensures that reported test results  
494 reflect true out-of-sample performance.



495  
496 **Figure 6.** Schematic diagram of temporal data partitioning. The chronological split ensures strict  
497 separation between training, validation, and test periods.  
498

#### 499 **Regression models**

500 Three regression models were evaluated for each selected predictor subset:

501 **Linear regression:** Ordinary least squares (OLS) regression, minimizing the sum of squared residuals.  
502 This model provides a simple baseline and interpretable coefficients, and was used during the  
503 predictor search phase due to its computational efficiency.

504 **LightGBM:** A gradient boosting framework using histogram-based learning, optimized for efficiency  
505 and scalability (Ke et al., 2017).

506 **XGBoost:** An optimized gradient boosting implementation with L1 and L2 regularization to prevent  
507 overfitting (Chen and Guestrin, 2016).

508 XGBoost consistently achieved the lowest mean MAE across all eight stations, and was therefore  
509 adopted as the primary model for reporting results. The following hyperparameters were used:  
510 learning rate  $\eta = 0.02$ , maximum tree depth  $d = 6$ , column subsampling ratio = 0.8, row subsampling  
511 ratio = 0.8, L1 regularization  $\alpha = 0.1$ , L2 regularization  $\lambda = 0.1$ . These values were selected based on  
512 common recommendations in the literature (Chen and Guestrin, 2016) rather than extensive tuning,  
513 as the focus of this study is predictor identification rather than model optimization.

#### 514 Predictor importance metrics

515 To analyze which predictors contribute most to model performance, XGBoost's built-in feature  
516 importance metrics were extracted. The *gain* metric was used as the primary importance measure.  
517 For a given predictor  $j$ , the gain  $G_j$  is defined as the sum of loss reductions achieved by all splits using  
518 that predictor:

$$519 \quad G_j = \sum_t \sum_{s \in S_{jt}} \Delta L(s)$$

520 where the outer sum is over all trees  $t$  in the ensemble,  $S_{jt}$  is the set of splits using predictor  $j$  in tree  
521  $t$ , and  $\Delta L(s)$  is the reduction in the loss function achieved by split  $s$ . This metric captures both the  
522 frequency with which a predictor is selected for splitting and the magnitude of improvement each  
523 split provides.

#### 524 Computational settings

525 The model was developed and tested on a standard virtual environment with 8 vCPUs (AMD EPYC  
526 9645) and 16 GB of RAM. This setup is intentionally equivalent to a standard consumer-grade  
527 desktop computer (e.g., an AMD Ryzen 7 system). Our goal was to ensure that the proposed  
528 forecasting method remains computationally accessible and can be implemented on affordable,  
529 everyday hardware without requiring expensive server clusters or specialized GPUs.

### 530 3.6 Evaluation Metrics

531 Model performance was assessed using multiple complementary metrics computed on the  
532 independent test set (2023–2024):

533 **Mean Absolute Error (MAE):** Primary performance metric, as mathematically defined above,  
534 representing the average magnitude of prediction errors in °C.

535 **Coefficient of determination ( $R^2$ ):** Proportion of variance in observed  $T_{min}$  explained by the model,  
536 defined as:

$$537 \quad R^2 = 1 - [\sum_i (y_i - \hat{y}_i)^2 / \sum_i (y_i - \bar{y})^2]$$

538 where  $\bar{y}$  is the mean of observed values.

539 **Bias:** Mean signed error, indicating systematic over- or under-prediction:

$$540 \quad Bias = (1/n) \times \sum_i (\hat{y}_i - y_i)$$

541 **Success rates:** Percentage of predictions within  $\pm 1$ °C and  $\pm 2$ °C of observations, providing  
542 operationally relevant accuracy measures.

543 **Error percentiles (P50, P90, P95):** Distribution of absolute errors, characterizing typical and worst-  
544 case performance.

545 **Seasonal and annual breakdowns:** MAE computed separately for each season (DJF, MAM, JJA, SON)  
546 and each year (2023, 2024) to identify temporal patterns in prediction difficulty.

547 Additionally, two baseline methods were computed for comparison: *persistence* (using the previous  
548 day's observed  $T_{min}$  as the forecast) and *climatology* (using the historical mean  $T_{min}$  for each  
549 calendar date). These baselines provide context for interpreting model skill (Jolliffe and Stephenson,  
550 2012).

551 **4. Results**

552 **4.1 Overall Performance**

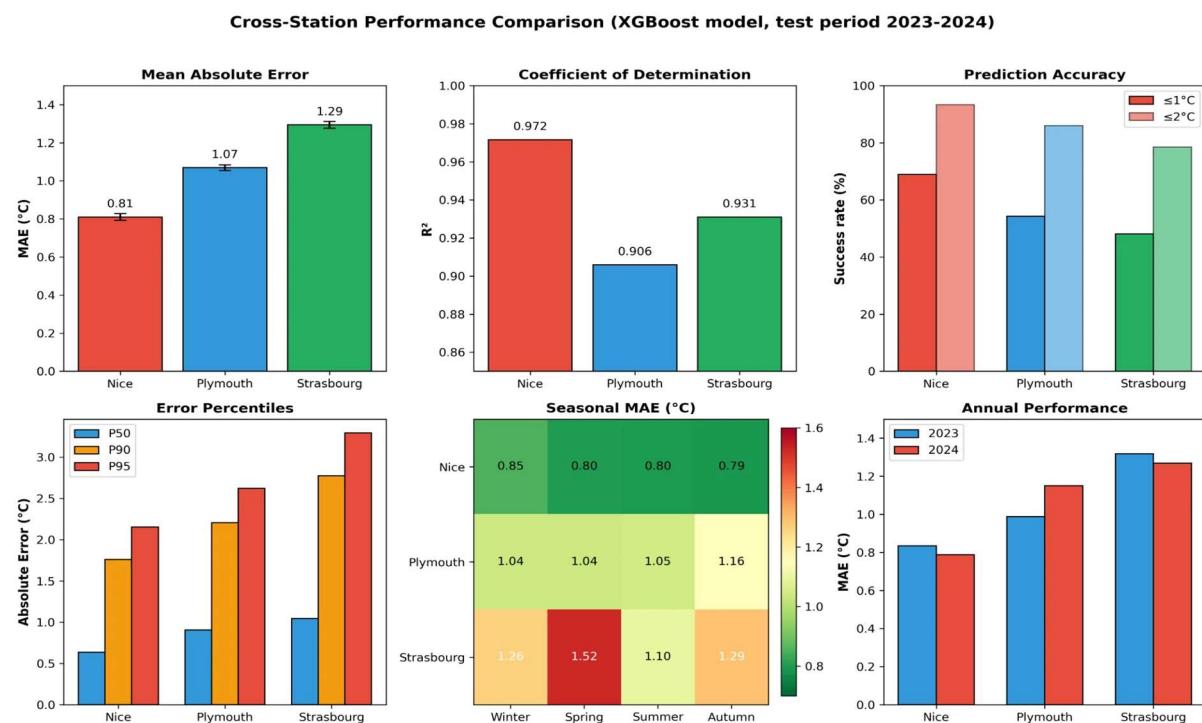
553 The predictor selection algorithm was applied to eight meteorological stations spanning diverse  
 554 European climates. Table 5 summarizes the prediction performance on the independent test period  
 555 (2023-2024) using the XGBoost regression model (Chen and Guestrin, 2016) driven by 90 selected  
 556 predictors.

557 **Table 5. Prediction performance for minimum temperature ( $T_{min}$ ) at eight European stations. MAE:**  
 558 **Mean Absolute Error; values are averages across five independent runs with different random seeds.**

Station	Climate	MAE (°C)	R <sup>2</sup>	Bias (°C)	≤1°C (%)	≤2°C (%)	P95 (°C)
Nice	Mediterranean	0.81	0.972	-0.13	68.9	93.3	2.15
Birmingham	Oceanic	0.92	0.936	-0.01	62.2	91.4	2.30
Plymouth	Oceanic	1.07	0.906	+0.00	54.3	86.0	2.62
Paris	Semi-continental	1.21	0.923	-0.35	50.1	82.1	3.00
Lyon	Semi-continental	1.22	0.946	-0.22	50.2	82.1	3.06
Edinburgh	Oceanic	1.24	0.892	-0.02	50.0	79.1	3.11
Strasbourg	Continental	1.29	0.931	-0.04	48.1	78.5	3.29
Brest	Oceanic	1.34	0.845	+0.11	45.7	77.1	3.31

559 Prediction accuracy varied substantially across stations, with MAE ranging from 0.81°C (Nice) to  
 560 1.34°C (Brest). Mediterranean climate (Nice) proved easiest to predict, with nearly 70% of forecasts  
 561 within 1°C of observations and over 93% within 2°C. Continental and oceanic climates presented  
 562 greater challenges, with success rates (≤2°C) between 77% and 91%, consistent with the greater  
 563 synoptic variability typically affecting these regimes (Wallace and Hobbs, 2006).

564 The 95th percentile of absolute errors (P95) provides insight into worst-case performance. Even for  
 565 the most challenging stations (Strasbourg, Brest), P95 remained below 3.5°C, indicating that large  
 566 errors were relatively rare.



568      **Figure 7.** Cross-station performance comparison for minimum temperature prediction (XGBoost model, test  
 569      period 2023-2024). Panels show mean absolute error, coefficient of determination, prediction accuracy, error  
 570      percentiles, seasonal MAE, and annual performance.

571

## 572      Comparison with baseline methods

573      To contextualize model performance, Table 6 compares results against two simple baselines: day-  
 574      ahead persistence (using yesterday's  $T_{min}$  as tomorrow's forecast) and climatology (historical mean  
 575       $T_{min}$  for the calendar date). Persistence is a well-known strong baseline for short-range temperature  
 576      forecasting (Wilks, 2011). A parsimonious SARIMAX model with annual Fourier terms (Box et al.,  
 577      2015) performs similarly to climatology and is clearly outperformed by persistence, confirming the  
 578      limited relevance of univariate linear models for daily minimum temperature prediction.

579      **Table 6.** Comparison of model performance against baseline methods. Improvement percentages  
 580      indicate MAE reduction relative to each baseline.

Station	Model	Persistence	Climatology	vs Pers.	vs Clim.
Nice	0.81	1.25	1.93	-35%	-58%
Birmingham	0.92	1.99	2.56	-54%	-64%
Plymouth	1.07	2.00	2.46	-46%	-56%
Paris	1.21	2.24	2.97	-46%	-59%
Lyon	1.22	2.25	2.98	-46%	-59%
Edinburgh	1.24	2.29	2.59	-46%	-52%
Strasbourg	1.29	2.42	3.09	-47%	-58%
Brest	1.34	2.40	2.73	-44%	-51%

581      The model reduced MAE by 35-54% compared to persistence and by 51-64% compared to  
 582      climatology. Notably, the improvement over persistence was smallest at Nice (35%), where the  
 583      Mediterranean climate naturally exhibits high day-to-day thermal stability. This finding is consistent  
 584      with the predictor analysis presented in Section 4.3, which reveals a strong persistence signal in the  
 585      selected predictors for this station.

## 586      Seasonal variations

587      Prediction difficulty varied seasonally, with distinct patterns emerging for different climate types  
 588      (Table 7).

589      **Table 7.** Seasonal MAE ( $^{\circ}$ C) for each station. Bold values indicate the most challenging season.

Station	Winter	Spring	Summer	Autumn	Most difficult
Nice	<b>0.85</b>	0.80	0.80	0.79	Winter
Birmingham	0.88	1.00	0.78	<b>1.02</b>	Autumn
Plymouth	1.04	1.04	1.05	<b>1.16</b>	Autumn
Paris	<b>1.33</b>	1.19	1.15	1.17	Winter
Lyon	<b>1.42</b>	1.19	1.10	1.18	Winter
Edinburgh	1.11	1.28	1.17	<b>1.40</b>	Autumn
Strasbourg	1.26	<b>1.52</b>	1.10	1.29	Spring
Brest	1.42	1.27	1.22	<b>1.46</b>	Autumn

590      Three distinct seasonal patterns emerged: (1) Oceanic stations (Birmingham, Plymouth, Edinburgh,  
 591      Brest) showed highest errors in autumn, likely due to increased synoptic variability during the  
 592      transition from summer to winter regimes (Wallace and Hobbs, 2006); (2) Semi-continental stations  
 593      (Paris, Lyon) performed worst in winter, when stable boundary-layer conditions and temperature  
 594      inversions frequently occur (Stull, 1988); (3) Strasbourg (continental) uniquely exhibited peak errors  
 595      in spring, when transitional weather between winter and summer patterns increases forecast  
 596      uncertainty.

597 **4.2 Predictor Analysis Methodology**

598 To investigate which atmospheric variables and spatial patterns drive minimum temperature  
599 predictions, detailed predictor analyses were conducted for three representative stations: Nice  
600 (Mediterranean, best performance), Plymouth (oceanic, intermediate), and Strasbourg (continental,  
601 most challenging). These stations were selected to maximize climatic diversity while spanning the  
602 full range of observed prediction accuracy.

603 **Variable-level versus grid-point stability**

604 A key methodological finding emerged from comparing predictor selections across five independent  
605 runs with different random seeds. While individual grid-point selections showed high variability  
606 (with no predictors appearing in all five runs for each station, except for the reference series itself),  
607 *variable-level* selections were remarkably stable: 23-24 of 26 ERA5 variables were consistently  
608 selected across all runs.

609 This apparent paradox reflects the high spatial autocorrelation of meteorological fields, whereby  
610 nearby grid points convey highly redundant information (von Storch and Zwiers, 1999). The  
611 algorithm identifies *which variables matter* robustly, even though the *exact spatial locations* vary  
612 between runs. Consequently, the following analyses focus on variable-level importance, aggregating  
613 contributions across all grid points of each variable.

614 **Importance metrics**

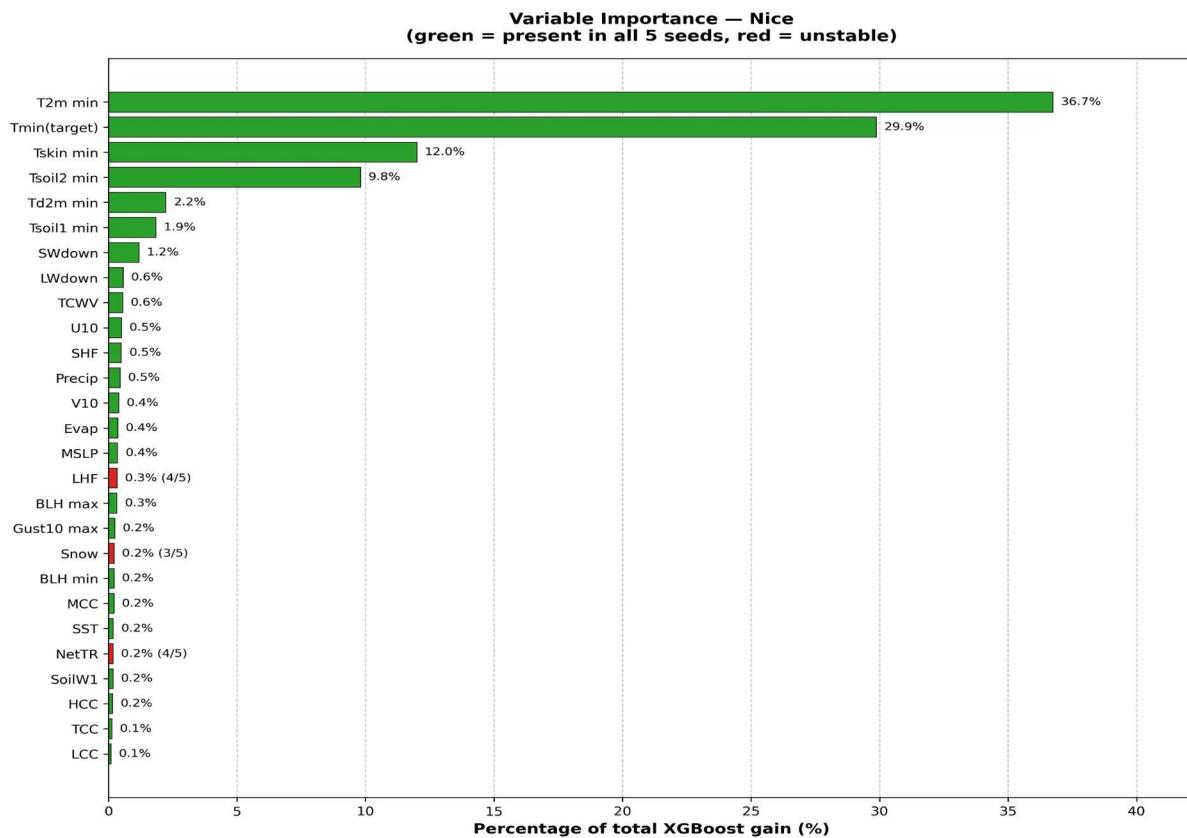
615 Predictor importance was quantified using XGBoost gain, which measures the total reduction in the  
616 loss function attributable to each predictor across all tree splits (Chen and Guestrin, 2016). This  
617 metric captures both the frequency with which a predictor is used and the magnitude of its  
618 contribution to model accuracy.

619 **4.3 Detailed Analysis: Nice (Mediterranean Climate)**

620 Nice, located on the French Riviera, exhibits a Mediterranean climate characterized by mild winters,  
621 warm summers, and thermal stability moderated by the sea. The station achieved the best  
622 prediction performance ( $MAE = 0.81 \pm 0.02^{\circ}\text{C}$ ,  $R^2 = 0.97$ ).

623 **Variable importance**

624 Figure 8 shows the relative importance of ERA5 variables for Nice. The predictor hierarchy revealed  
625 several distinctive features:



626

627 **Figure 8.** Variable importance for Nice (Mediterranean climate). Green bars indicate predictors  
 628 present in all 5 independent runs; red bars indicate unstable selections. The strong persistence signal  
 629 (Tmin target: 30%) is distinctive of this station.

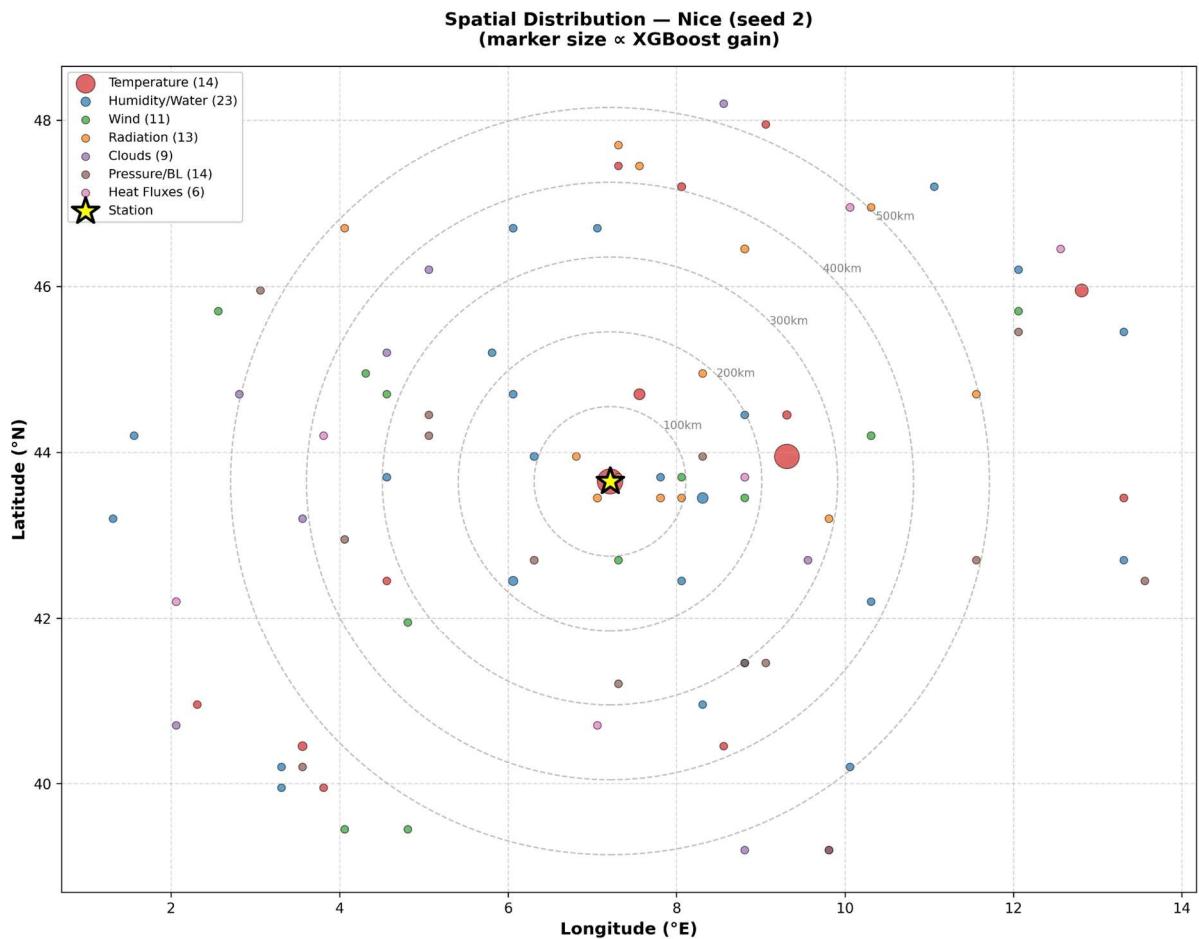
630 (1) 2-meter air temperature minimum (T2m\_min) dominated with 37% of total gain, though this  
 631 contribution was lower than at other stations (65-66% for Plymouth and Strasbourg). The dominant  
 632 role of near-surface temperature is consistent with the strong control exerted by local boundary-  
 633 layer processes under stable nocturnal conditions (Stull, 1988).

634 (2) The target variable itself (Tmin reference) contributed 30% of gain—a strong persistence signal  
 635 unique to Nice among the three analyzed stations. This reflects the thermal stability of  
 636 Mediterranean climate, where day-to-day temperature variations are buffered by maritime  
 637 influence. Such persistence effects are well documented in regions subject to strong oceanic thermal  
 638 inertia (Wilks, 2011; Oke, 1987).

639 (3) Surface temperatures (skin temperature 12%, soil temperature layer 2: 10%) were more  
 640 important than at other stations, capturing the land-sea thermal contrast that governs nocturnal  
 641 cooling in coastal Mediterranean environments (Oke, 1987).

#### 642 Spatial distribution

643 Nice exhibited the most spatially dispersed predictor pattern among the three stations, with only  
 644 65% of total gain originating within 200 km of the station (compared to 78% for Plymouth and  
 645 Strasbourg). Selected predictors extended across the northwestern Mediterranean basin, including  
 646 locations in Italy, southern France, and over the sea. This pattern suggests that the Mediterranean  
 647 basin functions as a coherent climatic unit for minimum temperature prediction, reflecting the  
 648 spatial coherence of synoptic-scale meteorological fields (von Storch and Zwiers, 1999).



649

650 **Figure 9.** Spatial distribution of selected predictors for Nice. Marker size is proportional to XGBoost  
651 predictors are dispersed across the northwestern Mediterranean basin, with only 65% of total  
652 gain within 200 km.

#### 653 4.4 Detailed Analysis: Plymouth (Oceanic Climate)

654 Plymouth, on the southwest coast of England, experiences a temperate oceanic climate  
655 characterized by mild temperatures year-round, high humidity, and frequent weather changes  
656 associated with Atlantic frontal systems. Prediction performance was intermediate (MAE =  $1.07 \pm$   
657  $0.01^\circ\text{C}$ ,  $R^2 = 0.91$ ).

#### 658 Variable importance

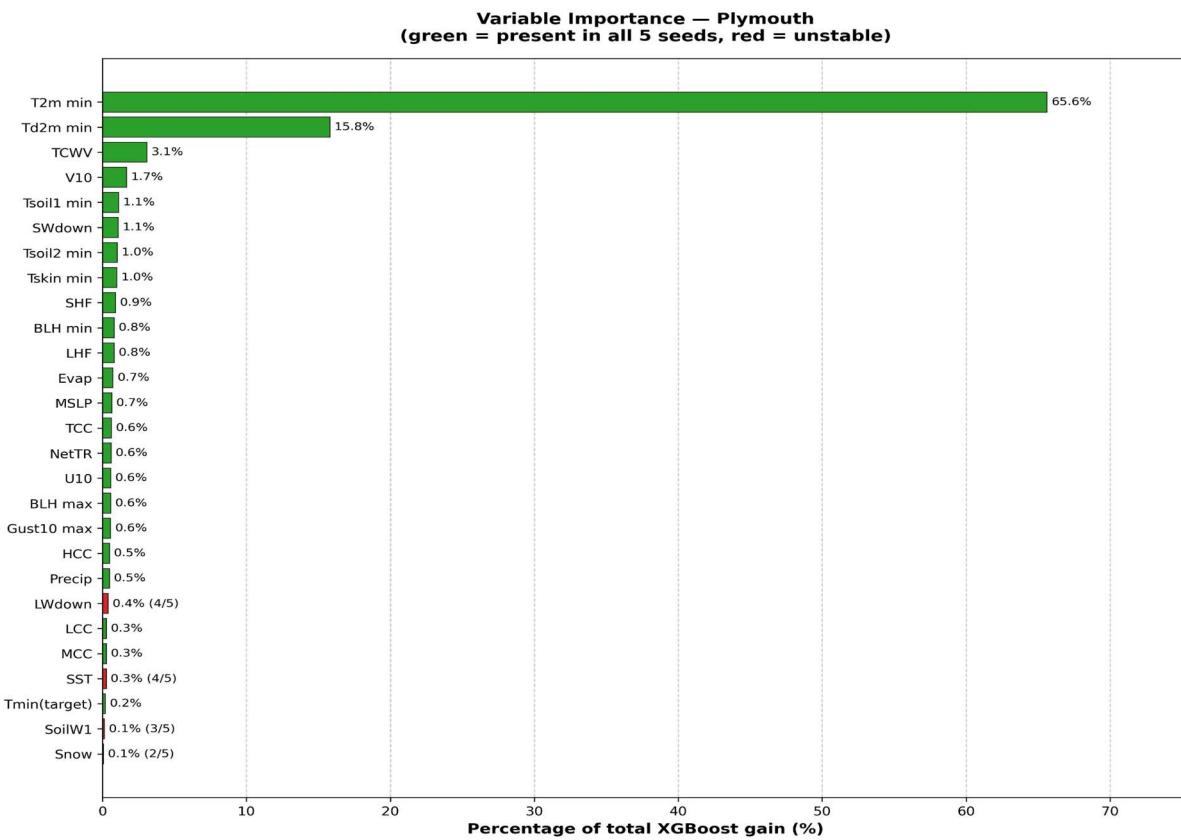
659 The predictor hierarchy for Plymouth differed markedly from Nice:

660 (1) *T2m\_min* strongly dominated with 66% of total gain, the highest concentration among the three  
661 stations.

662 (2) *Dewpoint temperature (Td2m)* contributed 16%—a distinctive signature of oceanic climate. The  
663 dewpoint captures humidity content of Atlantic air masses, which strongly influences nocturnal  
664 cooling rates through its effect on longwave radiation (Brutsaert, 1982).

665 (3) *Total column water vapour (TCWV: 3%) and meridional wind component (V10: 2%)* provided  
666 secondary contributions, reflecting the importance of moisture advection from the Atlantic.

667 (4) *The persistence signal was negligible* (*Tmin* reference  $< 1\%$ ), consistent with the changeable  
668 nature of oceanic weather where day-to-day conditions vary substantially under the influence of  
669 passing frontal systems (Wilks, 2011).



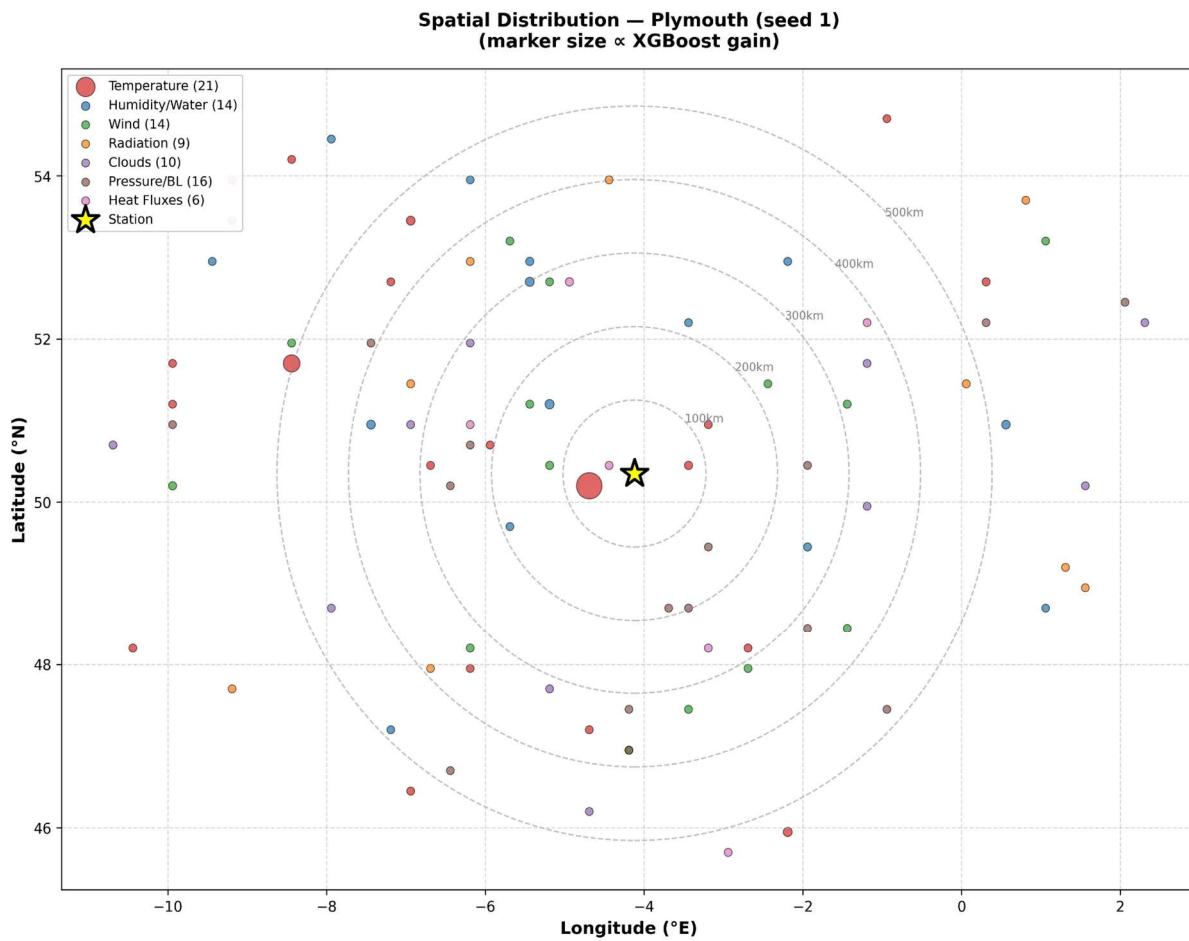
670

671 **Figure 10.** Variable importance for Plymouth (oceanic climate). T2m\_min strongly dominates (66%), with  
672 dewpoint temperature as distinctive oceanic signature.

673

#### 674 Spatial distribution

675 Predictor locations for Plymouth showed moderate spatial concentration, with 78% of gain within  
676 200 km. Selected grid points were distributed across southwest England, Wales, the English Channel,  
677 and Brittany, with humidity and wind predictors preferentially located in the Atlantic sector—  
678 consistent with the dominant westerly flow that characterizes this region's climate (Wallace and  
679 Hobbs, 2006).



680

681 **Figure 11.** Spatial distribution of selected predictors for Plymouth. Predictors extend across  
682 southwest England, Wales, and the English Channel, with humidity variables concentrated in the  
683 Atlantic sector.

684 Notably, Plymouth showed degraded performance in 2024 relative to 2023 (+0.16°C increase in  
685 MAE), the largest year-to-year variation among the eight stations. This may reflect unusual  
686 meteorological conditions during the test period that warrant further investigation.

#### 687 **4.5 Detailed Analysis: Strasbourg (Continental Climate)**

688 Strasbourg, located in the Upper Rhine Valley of northeastern France, exhibits a semi-continental  
689 climate with warm summers, cold winters, and large diurnal temperature ranges. This station  
690 presented the greatest prediction challenge among the three analyzed (MAE =  $1.29 \pm 0.02^\circ\text{C}$ ,  $R^2 =$   
691 0.93).

#### 692 **Variable importance**

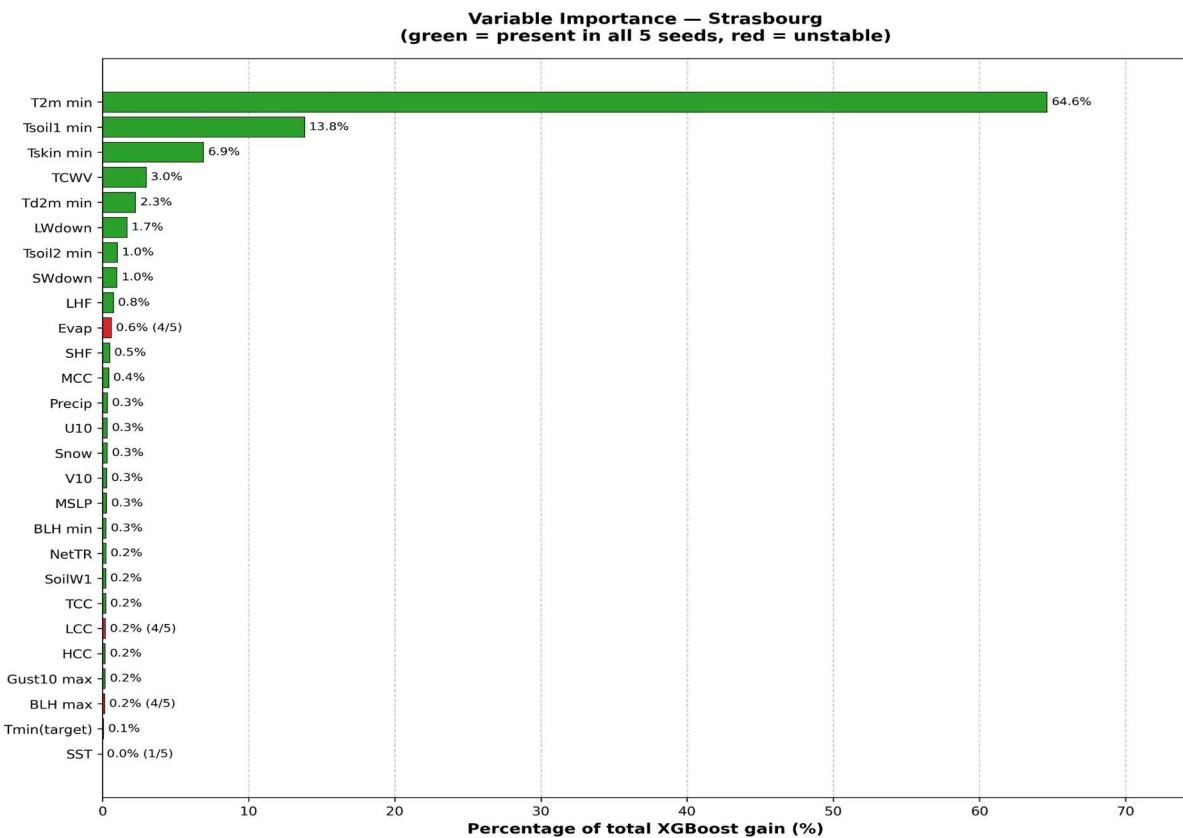
693 The predictor hierarchy for Strasbourg featured:

694 (1) *T2m\_min* dominated with 65% of total gain, similar to Plymouth.

695 (2) *Soil temperature (layer 1)* contributed 14%—a distinctive continental signature absent from the  
696 oceanic and Mediterranean stations. This reflects the role of ground thermal inertia in modulating  
697 nocturnal cooling under the clear skies typical of continental high-pressure systems (Oke, 1987).

698 (3) *Skin temperature* (7%) and *total column water vapour* (3%) provided secondary contributions.

699 (4) *Persistence was minimal* (*Tmin* reference < 1%), indicating that continental climate, like oceanic  
700 climate, exhibits substantial day-to-day variability in minimum temperatures (Wilks, 2011).

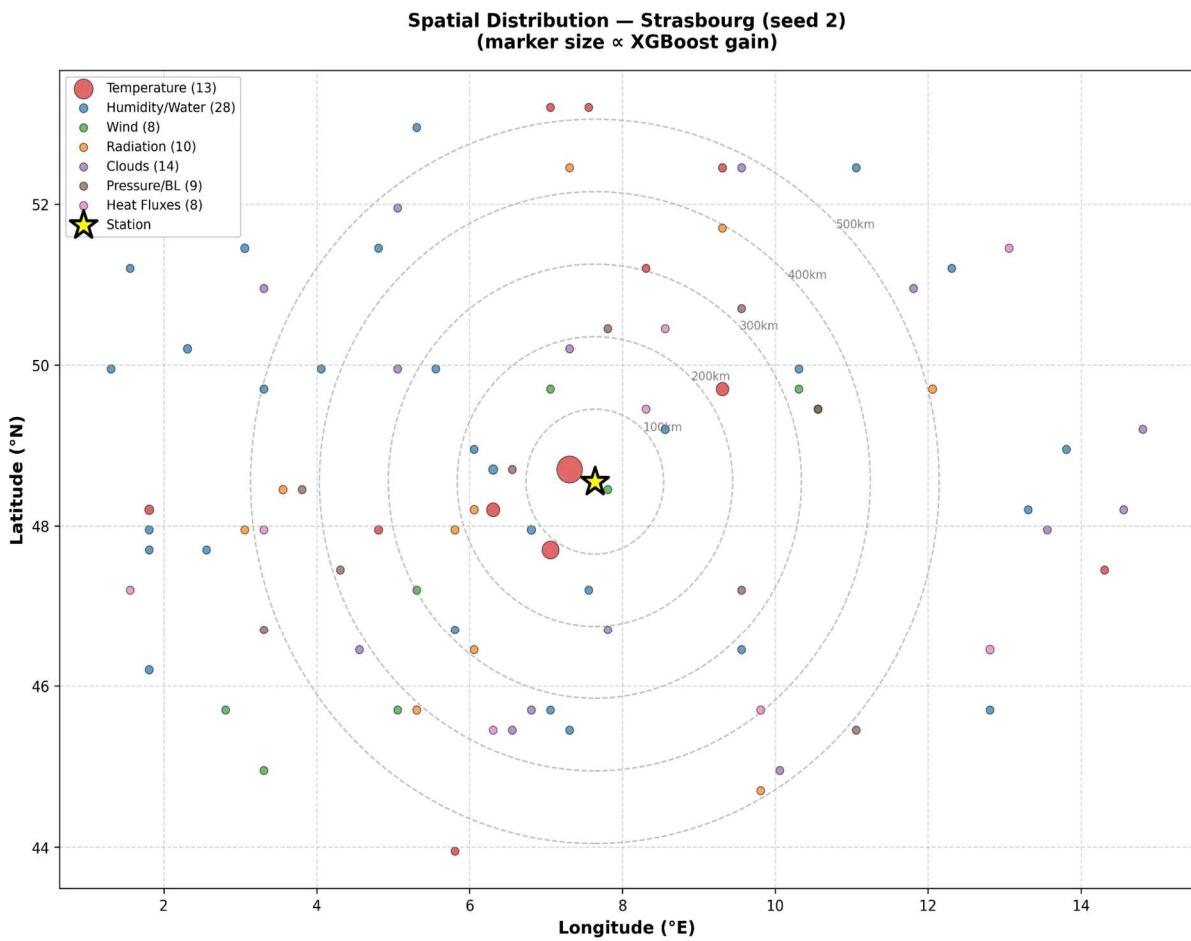


701

702 **Figure 12.** Variable importance for Strasbourg (continental climate). Soil temperature (layer 1: 14%)  
703 emerges as a distinctive continental signature.

#### 704 Spatial distribution and gain concentration

705 Predictor locations concentrated in the Rhine Valley and surrounding regions, with 78% of gain  
706 within 200 km of the station. Some predictors were located in western France, potentially capturing  
707 the influence of Atlantic air masses that occasionally penetrate into the continental interior (Wallace  
708 and Hobbs, 2006).



709

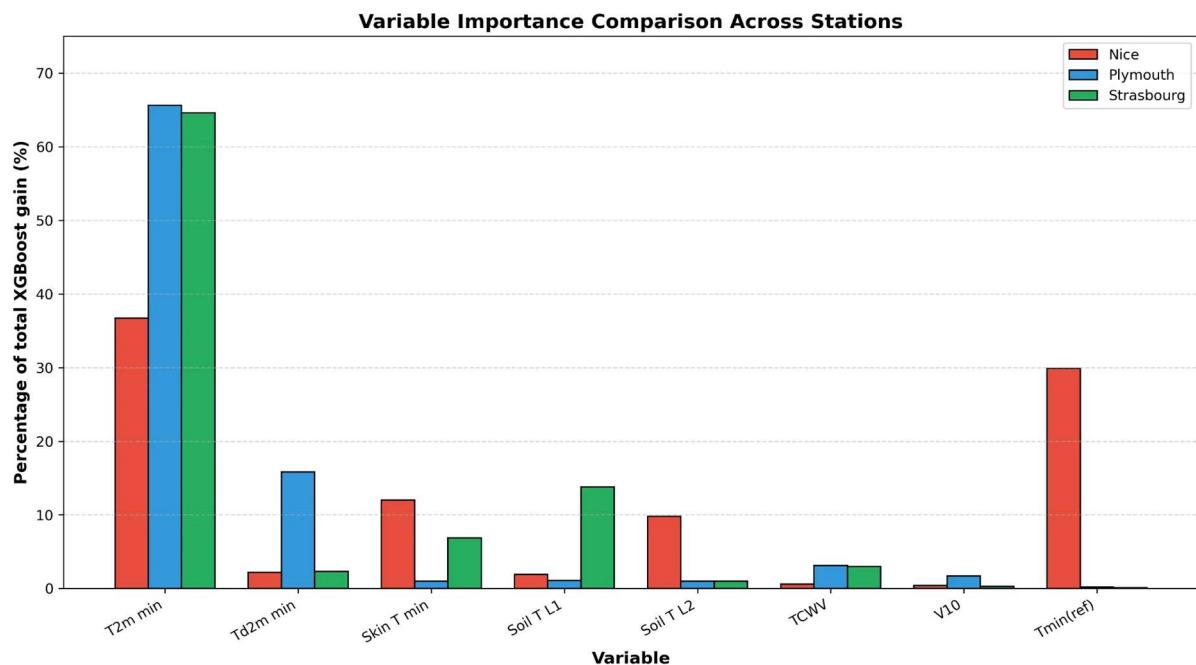
710 **Figure 13.** Spatial distribution of selected predictors for Strasbourg. Predictors concentrate in the  
711 Rhine Valley region, with 78% of gain within 200 km of the station.

712 Analysis of gain concentration revealed extreme inequality in predictor contributions: the top 4  
713 predictors accounted for 80% of total gain, and the top 12 for 90%. The remaining 78 predictors  
714 contributed marginally, suggesting that a substantially reduced predictor set might achieve  
715 comparable performance.

716 Spring presented the greatest prediction challenge for Strasbourg (MAE = 1.52°C), unique among the  
717 analyzed stations. This likely reflects the transitional character of spring weather in continental  
718 climates, where rapidly shifting patterns between residual winter conditions and emerging summer  
719 regimes create high forecast uncertainty (Wallace and Hobbs, 2006).

## 720 [4.6 Cross-Station Synthesis](#)

721 General patterns



722  
723 **Figure 14.** Variable importance comparison across the three representative stations. *T2m\_min* dominates at all  
724 stations (37-66%), with climate-specific secondary predictors.

725  
726 Despite substantial differences in climate and prediction difficulty, several patterns emerged  
727 consistently across all three stations:  
728 (1) *Near-surface air temperature (T2m\_min)* was universally the dominant predictor, contributing  
729 37-66% of total gain.  
730 (2) *The top 3 variables captured 85-90% of total predictive information*, indicating that minimum  
731 temperature forecasting is fundamentally driven by a small set of thermal variables, in line with  
732 established physical understanding of nocturnal cooling processes (Stull, 1988).  
733 (3) *Predictive information was concentrated locally*, with 65-78% of gain originating within 200 km of  
734 each station, consistent with the spatial coherence of synoptic and mesoscale meteorological fields  
735 (von Storch and Zwiers, 1999).  
736 (4) *Variable-level selection was highly stable* (23-24 of 26 variables present in all five independent  
737 runs), while grid-point selection showed high variability due to spatial autocorrelation, a known  
738 characteristic of gridded atmospheric data (von Storch and Zwiers, 1999).

#### 739 Climate-specific signatures

740 Beyond these universal patterns, each climate type exhibited distinctive predictor signatures (Table  
741 8).

742 **Table 8.** Climate-specific predictor signatures identified through the selection algorithm.

Climate	Distinctive predictor	Contribution	Physical interpretation
Mediterranean	Tmin reference	30%	Thermal stability (maritime buffer)
Oceanic	Dewpoint (Td2m)	16%	Atlantic humidity advection
Continental	Soil temperature	14%	Ground thermal inertia

743 These signatures align with physical understanding of each climate regime. Mediterranean climates  
744 showed strong persistence linked to maritime buffering (Oke, 1987), oceanic climates were  
745 characterized by the importance of humidity-related predictors associated with Atlantic air masses  
746 (Brutsaert, 1982), and continental climates exhibited a clear influence of soil thermal inertia (Oke,

747 These signatures are physically consistent and reinforce the interpretability of the predictor  
748 selection methodology.

#### 749 **Implications for predictor selection**

750 The extreme concentration of predictive gain (80% from top 4 predictors, 90% from top 12) suggests  
751 that the 90-predictor configuration may be over-specified for operational applications. A reduced set  
752 of 20-30 predictors, focused on thermal variables and climate-specific secondary predictors, would  
753 likely achieve comparable performance with reduced computational cost. However, retaining the  
754 broader predictor pool in the research context proved valuable for identifying the climate-specific  
755 signatures documented above.

## 756 **5. Conclusion**

### 757 **5.1 Summary of findings**

758 This study developed and validated an automated framework for systematic spatial exploration and  
759 predictor selection applied to next-day minimum temperature forecasting across Western Europe. By  
760 combining ERA5 reanalysis data with gradient boosting regression techniques, we identified optimal  
761 predictor configurations for eight climatically diverse reference sites, achieving substantial  
762 improvements over baseline methods.

763 The results demonstrate that minimum temperature forecasting benefits significantly from spatially-  
764 distributed predictors, with model performance reducing MAE by 35–54% compared to persistence  
765 and 51–64% compared to climatological means. Mediterranean climates proved easiest to predict  
766 (MAE = 0.81°C), while oceanic and continental sites presented greater challenges (MAE = 1.07–1.34°C).  
767 These performance differences reflect fundamental distinctions in the physical processes governing  
768 nocturnal cooling across climate regimes.

769 A key finding concerns the emergence of climate-specific predictor signatures. Mediterranean stations  
770 exhibited strong persistence signals driven by maritime thermal buffering, oceanic climates showed  
771 enhanced sensitivity to dewpoint temperature reflecting Atlantic humidity advection, and continental  
772 sites featured significant soil temperature contributions linked to ground thermal inertia. These  
773 signatures align with established meteorological understanding and demonstrate that the search  
774 algorithm successfully identifies physically meaningful predictors.

775 The extreme concentration of predictive gain—with 80% originating from just 4 predictors and 90%  
776 from 12—suggests that operationally efficient configurations could be derived from our  
777 comprehensive predictor analysis. The single-day temporal window proved optimal for 75% of stations  
778 tested, consistent with the principle of parsimony in time series modeling.

### 779 **5.2 Broader perspectives and future applications**

780 This work is part of a broader project aimed at leveraging time series analysis for predictive  
781 applications across multiple domains. The methodology presented here—systematic spatial  
782 exploration combined with machine learning-based predictor selection—is inherently generic and can  
783 be adapted to various forecasting challenges beyond minimum temperature prediction.

784 Preliminary experiments on river gauge level forecasting (e.g., the Mississippi River in St. Louis,  
785 Missouri) have yielded promising results, with the algorithm successfully identifying informative  
786 predictors using temporal window sizes larger than one day (e.g., D-3).

787 This application demonstrates the framework's potential for hydrological forecasting, where accurate  
788 water level predictions are essential for flood management and water resource planning. Other  
789 potential applications span diverse fields, including air quality forecasting (pollutant concentration

790 prediction), agricultural meteorology (frost risk assessment), energy sector planning (demand  
791 forecasting), and domains entirely unrelated to Earth sciences, using dimensionalities beyond the  
792 geographic one.

793 While the present study used daily temporal resolution, the algorithm can be readily adapted to other  
794 time units appropriate to the problem at hand—hours for short-term operational forecasting, weeks  
795 or months for seasonal predictions, or even years for long-term trend analysis (e.g., as in seasonal  
796 forecasting frameworks discussed by Hyndman and Athanasopoulos, 2021). This temporal flexibility,  
797 combined with the spatial exploration capabilities demonstrated here, positions the framework as a  
798 versatile tool for time series prediction across scales and domains, aligning with broader applications  
799 of machine learning in environmental sciences (Reichstein et al., 2019).

### 800 5.3 Operational considerations

801 While the predictor selection algorithm explores a vast search space (~45,000 candidates) to identify  
802 the optimal 90-predictor set per site, it is designed for infrequent execution—typically once per year  
803 or every two years—to incorporate newly available data and maintain relevance under evolving  
804 climate conditions. This process requires approximately one hour of computation per site on a  
805 standard personal computer, making it feasible for periodic updates without significant resource  
806 demands.

807 In contrast, daily forecasting operations leverage the pre-selected predictor set: the XGBoost  
808 regression model is applied using current-day values from these 90 predictors, completing inference  
809 in mere tens of seconds. This separation between offline optimization (guided by MAE as a directional  
810 metric) and online prediction ensures scalability and efficiency, allowing the framework to be  
811 deployed in real-time systems while focusing primarily on discovering robust, physically meaningful  
812 predictor configurations rather than minimizing absolute error in isolation, as emphasized in  
813 operational ML guidelines (Sculley et al., 2015).

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### 883 **Code availability**

884 All code developed for this study—including the predictor selection algorithm, data processing  
885 pipelines, analysis scripts, model training configurations, and results files (json)—will be made  
886 publicly available upon publication via repositories such as GitHub and Zenodo, with a permanent  
887 DOI assigned to the archive. The exact links will be provided upon request. A Cython-compiled high-  
888 performance version of the selection algorithm will also be provided. Docker container ensuring full  
889 reproducibility of the computational environment will be released alongside the code repository.  
890 This repository will be publicly released within 8 to 10 weeks of this preprint.

### 891 **Data availability**

892 ERA5 reanalysis data are freely available from the Copernicus Climate Data Store  
893 (<https://cds.climate.copernicus.eu>). Daily minimum temperature observations were obtained from  
894 NOAA's National Centers for Environmental Information Global Summary of the Day dataset  
895 (<https://www.ncei.noaa.gov>).

### 896 **Author contribution**

897 Eric Duhamel designed the study, developed the methodology, wrote the code, performed the  
898 analyses, and prepared the manuscript.

### 899 **Competing interests**

900 The author declares that there is no conflict of interest.

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