

1 **GRAIN FRICTION CONTROLS**
2 **CHARACTERISTICS OF SEISMIC CYCLE IN**
3 **ROUGH FAULTS WITH GRANULAR GOUGE**

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24 **Abstract:**

25 Mature faults with rough surface at their core contain granular gouge, however, the
26 presence of fault gouge is mostly disregarded in the analysis of effect of fault's surface
27 roughness on the mechanics and stability of faults. In this work, we consider a rough
28 fault system with constant roughness at the wall-gouge interaction and study the
29 effect of grain friction on the characteristics of seismic cycles. Our discrete element
30 simulations show that the stick-slip frictional strength and dilation of the rough fault
31 system, as well as their variations, nonlinearly increase with the particle friction, but
32 at high particle friction saturate. By statistical analyses on a large number of slip
33 events, we find that the average recurrence time and its variations decrease with
34 particle friction. A rough fault with higher grain friction shows more small slip
35 events, but also contains a limited number of extreme events and demonstrates a
36 more complex nucleation phase with higher stored energy. We analyze the pseudo
37 acoustic emission, which is based on monitoring of the velocity signal of particles,
38 and find higher temporal and more spatially distributed acoustic emissions for rough
39 fault with higher grain friction. Our findings in this study show that, in rough faults
40 with granular gouge, where the fault zone walls are totally engaged to the granular
41 gouge, the friction at grain scale controls the characteristics of stick-slip cycles
42 showing similar influence on the mechanics of faults as the roughness of fault's
43 surface in absence of fault gouge.

44 **Keywords:** friction, roughness, stick-slip, granular materials, fault gouge, fault
45 mechanics

46 **1- Introduction**

47 Mature faults at their core are usually characterized by a granular gouge layer, created
48 by wear, comminution and other frictional processes in nature [Engelder, 1974;
49 Shimamoto, 1979; Chester et al., 1985; Chester & Logan, 1986; Marone et al., 1990;
50 Chester & Chester, 1998; S. Cashman & Cashman, 2000; Chester & Chester, 2000;
51 Faulkner et al., 2003; Heermance et al., 2003; S. M. Cashman et al., 2007; Zoback et
52 al., 2010]. Characteristics of slip events in laboratory stick-slip experiments including
53 stress drop, pre-seismic friction, recurrence time, compaction etc. have been studied
54 widely for different loading configurations and particle properties [Dieterich &
55 Kilgore, 1994; Chris Marone, 1998; Karner & Marone, 2000; Mair et al., 2002;
56 Brantut et al., 2008; Rathbun & Marone, 2010; Haines et al., 2014; Rosenau et al.,
57 2017]. It is shown that frictional strength and stability of a sheared granular zone are
58 sensitive to grain shape, particle size distribution and their evolution [Mair et al.,
59 2002; Anthony & Marone, 2005].

60 The stability and slip behavior of natural faults and their relation with fault
61 properties, and in particular their geometry and surface roughness, have been under
62 debate during last years. The fault roughness spans from microns to tens of
63 kilometers [Kozłowska et al., 2018] and can be measured using different methods
64 [Renard et al., 2006; Sagy et al., 2007; Candela et al., 2009; Bistacchi et al., 2011; Emily
65 E. Brodsky et al., 2011; Candela et al., 2012; Emily E Brodsky et al., 2016]. At field
66 scale, the fault roughness is suggested to control the stress drop and slip distribution
67 during earthquakes, hydraulic fracturing and subduction of seafloor relief [Bouchon

68 et al., 2010; Candela, Renard, Bouchon, et al., 2011; Candela, Renard, Schmittbuhl,
69 et al., 2011; Wang & Bilek, 2014; Rijnsingen et al., 2018]. At lab scale, experimental
70 studies showed that the topology and fault structure influence the spatial and
71 temporal distribution of small and large earthquakes [Ohnaka & Shen, 1999; Ohnaka,
72 2003; T. H. W. Goebel, Becker, et al., 2014; Thomas H.W. Goebel et al., 2017].
73 Numerical simulations have been also used to study the effect of fault roughness on
74 mechanics and nucleation of slip events [Chester & Chester, 2000; Dieterich &
75 Smith, 2009; Angheluta et al., 2011; Dunham et al., 2011; Fournier & Morgan, 2012;
76 Rathbun et al., 2013; Bruhat et al., 2016; Zielke et al., 2017; Tal & Hager, 2018; Tal
77 et al., 2018].

78 However, the presence of granular fault gouge is neglected in most of the
79 previous studies related to fault roughness. The study of Rathbun et al. [2013],
80 considered the presence of fault gouge during stable sliding simulations and found
81 that, on the first-order, fault strength is controlled by particle friction and mechanical
82 coupling of the fault zone wall to the gouge for both rough and smooth faults.
83 Rathbun et al. [2013] showed that, for rough faults, when the fault zone walls and
84 gouge zone are totally engaged with no slip between them, only the gouge friction
85 controls the strength. In this work, we systematically vary the particle friction
86 coefficient (called here for simplicity particle friction) in a fault with granular gouge
87 and focus on stick-slip dynamics to study the effect of particle friction on the slip
88 size distribution and inter-event time of seismic cycles describing the
89 micromechanical properties of frictional processes that take place in the fault damage

90 zone. To this end, we perform 3-D DEM simulations recording hundreds of slip
 91 events for statistical analyses provided by the advantage of numerical simulations.
 92 We also study the evolution of elastic strain energy (i.e. potential energy in
 93 [Dorostkar, 2018; Dorostkar & Carmeliet, 2018]), kinetic energy, and the
 94 micromechanics of slip events and explain the macro-scale response of fault with
 95 grain-scale metrics (e.g. [Dorostkar, Guyer, et al., 2017a]). Using a simple rheological
 96 and geometrical granular contact model, we will show that, although our simplified
 97 numerical approach does not contain all complexities present in nature, it shows that
 98 for faults with granular gouge, the friction of the gouge particles controls the
 99 characteristics of stick-slip cycles. We also compare our findings with previous
 100 numerical and experimental works and show that, in rough faults with granular
 101 gouge, the friction at grain scale has similar influence on the mechanics of faults as
 102 the roughness at fault's surface in absence of fault gouge.

103 **2- Model description**

104 Due to particulate nature of granular fault gouge, we use DEM to model gouge
 105 grains. In DEM, the equations of motion solved for each particle considering the
 106 applied forces are:

$$107 \quad \sum F_p = m \left(\frac{d}{dt} u_p \right), \quad (1)$$

$$108 \quad \sum T_p = I \left(\frac{d}{dt} \omega_p \right), \quad (2)$$

109 where m , I , u_p and ω_p are the mass, the moment of inertia and the translational and
 110 angular velocity of particle, respectively. In Eq. (1) and (2), F_p and T_p are the forces

111 and torques acting on particle i owing to particle-particle contacts. In soft sphere
112 DEM, the particle-particle contact allows an overlap between them and the contact
113 law is described by a combination of different rheological elements (spring, dashpot,
114 slider etc.). These elements are at play when particles are in contact. Upon contact
115 loss, there are no contact forces acting between the particles. In order to capture
116 particle scale nonlinearity, we use the nonlinear Hertzian contact law. In this particle-
117 particle contact law, the spring stiffness and the coefficient of damping are function
118 of particle material properties and the overlap between particles [Hertz, 1882; Di
119 Renzo & Di Maio, 2004]. The normal and tangential contact forces are calculated as
120 follows:

$$121 \quad F_{pn} = -k_{pn}\delta\varepsilon_{pn} + c_{pn}\delta u_{pn}, \quad (3)$$

$$122 \quad F_{pt} = \min \left\{ \left| k_{pt} \int_{t_{c,0}}^t \delta u_{pt} dt + c_{pt} \delta u_{pt} \right|, \mu_c F_{pn} \right\}, \quad (4)$$

123 where k_{pn} and k_{pt} are the normal and tangential spring stiffness, c_{pn} and c_{pt} are
124 the normal and tangential damping coefficient, $\delta\varepsilon_{pn}$ is the overlap and δu_{pn} and
125 δu_{pt} are the relative normal and tangential velocities of two particles in contact,
126 respectively. In Eq. (4), the parameter μ_c represents the particle friction coefficient
127 that limits the tangential force. When the tangential contact force between two
128 particles in contact reaches this limit, they start sliding against each other. The integral
129 term in Eq. (4) shows an incremental spring, storing energy based on the relative
130 elastic tangential deformation of the particle surface starting from the moment
131 particles touch each other at $t_{c,0}$. A damping is added to the spring component of

132 the tangential force if the Coulomb criterion is not met [Di Renzo & Di Maio, 2004;
 133 Goniva et al., 2012]. The spring and damping coefficients are calculated as follows:

$$134 \quad k_{pn} = \frac{4}{3} Y^* \sqrt{R^* \delta \varepsilon_{pn}}, \quad (5)$$

$$135 \quad k_{pt} = 8 G^* \sqrt{R^* \delta \varepsilon_{pn}}, \quad (6)$$

$$136 \quad c_{pn} = -2 \sqrt{\frac{5}{6}} \times \frac{\ln(r)}{\sqrt{\ln^2(r) + \pi^2}} \times \sqrt{2Y^* \sqrt{R^* \delta \varepsilon_{pn}} m^*}, \quad (7)$$

$$137 \quad c_{pt} = -2 \sqrt{\frac{5}{6}} \times \frac{\ln(r)}{\sqrt{\ln^2(r) + \pi^2}} \times \sqrt{8 G^* \sqrt{R^* \delta \varepsilon_{pn}} m^*}, \quad (8)$$

138 where, r is the restitution coefficient, and Y^* , R^* , G^* and m^* are the equivalent
 139 Young's modulus, radius, shear modulus and mass, respectively, calculated as
 140 follows:

$$141 \quad \frac{1}{Y^*} = \frac{(1-\nu_1^2)}{Y_1} + \frac{(1-\nu_2^2)}{Y_2}, \quad (9)$$

$$142 \quad \frac{1}{G^*} = \frac{2(2-\nu_1)(1+\nu_1)}{Y_1} + \frac{2(2-\nu_2)(1+\nu_2)}{Y_2}, \quad (10)$$

$$143 \quad \frac{1}{R^*} = \frac{1}{R_1} + \frac{1}{R_2}, \quad (11)$$

$$144 \quad \frac{1}{m^*} = \frac{1}{m_1} + \frac{1}{m_2}, \quad (12)$$

145 where subscripts 1 and 2 refer to the two particles in contact and ν is the Poisson's
 146 ratio of the particle.

147 Figure 1 illustrates a granular layer representing a part of a granular fault gouge.
 148 In our model, 8000 spherical particles constitute this layer with particle diameter
 149 ranging 90-150 μm having a uniform, poly-disperse particle size distribution. The
 150 sample size in our simulations is $11 \times 1.5 \times 0.8 \text{ mm}^3$. In order to model a fault with

151 rough surface in a simplified way, on the sample top and bottom, we employ two
152 corrugated plates with high surface roughness modeled by a friction coefficient of
153 0.9 between the plates and particles to facilitate the transmission of shear stresses to
154 the granular gouge (see Fig.1 insets). This kind of geometry is inspired from the
155 BIAX laboratory earthquake machine using corrugated driving blocks [Chris
156 Marone, 1998; Rivière et al., 2018] and simulates a fault system with rough surface.
157 Our analysis shows that use of these corrugated plates with high surface friction leads
158 to a total engagement of fault blocks with granular gouge. Therefore, no slip takes
159 place at the boundary between the plates and the granular gouge, but the deformation
160 is distributed across the sample thickness, inside the granular gouge (see Fig. 13). In
161 this paper, our study is dedicated to the influence of the gouge particle friction on
162 stick-slip behavior and we do not study the effect of fault roughness i.e.
163 characteristics of the corrugated plates e.g. length, depth, etc. on the dynamics of the
164 sheared granular layer. A detailed analysis showed that change of these characteristics
165 does not fundamentally alter the dynamic regime of stick-slip. Rathbun et al. [2013]
166 showed that, due to a better engagement between the corrugated plates and gouge
167 zone, a higher roughness of corrugated plates characterized with larger tooth size will
168 extend the deformation more inside the gouge zone [Rathbun et al., 2013]. The
169 previous research showed that by increasing the rolling friction i.e. the resistance of
170 particles to rotate, the stick-slip dynamics becomes more stabilized and the dynamic
171 regime tends to change from stick-slip to slow slip, defined as a more gradual
172 decrease of the macroscopic friction rather than a sudden drop [Dorostkar, 2018].

173 Therefore, since our focus here is on stick-slip dynamics and we aim to isolate the
174 effect of particle friction in a detailed analysis, we do not include the effect of rolling
175 resistance and consider it zero within the scope of this manuscript.

176 On the front- and back-side of the sample, we implement frictionless walls with
177 the same elastic properties of particles. This type of interaction between particles and
178 walls are designed to avoid rigid wall boundary conditions. Wall-particle interaction
179 in our DEM model is the same as particle-particle interaction when one particle has
180 an infinite radius. Periodic boundary conditions are applied at the left and right
181 sidewalls representing a long fault gouge in x direction. The periodic boundary
182 conditions allow for large shear displacements, and facilitate recording many slip
183 events to be used for statistical analyses. To prepare the sample, particles are inserted
184 randomly in space descending with an initial velocity of 10^{-2} cm/s. Next, the upper
185 plate is moved downward to apply a confining stress to confine the sample (applied
186 on top x - y plane in Fig.1). At this stage, the load increases until the desired confining
187 stress is attained (10 MPa). The position of the upper plate is adapted continuously,
188 as in the lab experiments, in order to maintain the confining stress constant. At
189 constant confining stress, shearing is initiated by moving the bottom plate in x
190 direction with a displacement-controlled mechanism (constant velocity of $600 \mu\text{m/s}$)
191 until reaching the maximum shear stress, at which point the stick-slip process
192 commences. The particle density is 2900 kg/m^3 that results in an applied time step
193 of 15×10^{-9} seconds for DEM calculations, within the recommended range based on
194 the Rayleigh time. Our DEM calculations remain in the quasi-static regime by

195 controlling the inertial number to be below 10^{-3} [MiDi, 2004; Sheng et al., 2004;
196 Agnolin & Roux, 2007]. Similar to our previous works [Dorostkar, Johnson, et al.,
197 2017; Dorostkar et al., 2018], we use LIGGGHTS [Goniva et al., 2012; Kloss et al.,
198 2012] to model the granular fault gouge.

199 **3- Results**

200 We show in Fig. 2a and 2b the evolution of macroscopic friction and gouge thickness
201 during the stick-slip dynamics, respectively, for three different particle friction values
202 of 0.1, 0.5 and 0.9. The macroscopic friction is defined as the ratio of shear stress to
203 normal stress (confining stress) on the driving block and the gouge thickness
204 represents the measurement in z direction of Fig. 1. All simulations start from the
205 same initial thickness and particle arrangement. Figure 2c shows with higher
206 resolution the shaded area in Fig. 2a. The macroscopic friction rapidly increases in
207 the sample and after some stable sliding; the sample undergoes stick-slip dynamics,
208 where the start of stick-slip dynamics is denoted by time zero. For each stick-slip
209 cycle, the macroscopic friction (or shear stress, since the confining stress is kept
210 constant) increases nonlinearly reaching to a critical state where micro-slips take place
211 followed by a major slip event. The average macroscopic friction increases
212 nonlinearly with the particle friction: at low particle friction values (0.1 to 0.5), the
213 macroscopic friction increases much more compared to high particle friction values
214 (0.5 and 0.9). A similar behavior can be observed for the gouge thickness (Fig. 2b).

215 The schematic stick-slip cycle shown in Fig. 3a shows the definition of recurrence
216 time, micro-slips and major slip event in our analysis. We recognize major slip events

217 with a drop in macroscopic friction larger than or equal to 0.01, a threshold that
218 avoids capturing micro-slips before a major slip event [Dorostkar, Guyer, et al.,
219 2017b]. We perform long simulations and measure the average macroscopic friction,
220 gouge thickness and slip recurrence time over all stick-slip phases as function of
221 particle friction (Fig. 3b-d). We observe that the gouge strength represented by
222 macroscopic friction increases nonlinearly with particle friction and saturates at
223 around $\mu_c = 0.9$ to 1. A similar behavior is observed for gouge thickness (Fig. 3c).
224 The standard deviations of macroscopic friction signal and gouge thickness are found
225 to increase with increasing particle friction (Fig. 3b and 3c). The slip recurrence time
226 and its standard deviation decreases with increasing particle friction, meaning that
227 slip events occur more often and more regularly in a gouge with higher frictional
228 particles (Fig. 3d)

229 To complement the observations in Fig. 3, in Fig. 4a, we observe an almost linear
230 relation between macroscopic friction and gouge thickness for simulations with
231 different particle friction. The more dense population of data points at higher
232 thickness and macroscopic friction is consistent with the nonlinear behavior
233 observed in Figs. 2b and 2c. We also observe that the number of slip events increases
234 with increasing particle friction (equivalent to a decreasing recurrence time as seen in
235 figure 3d). The histogram of slip events' friction drop (Fig. 5) shows that this increase
236 in number of events at higher particle frictions mainly stems from smaller events.
237 This implies that, a fault gouge with higher particle friction experiences more smaller
238 slip events at a shorter and more regular inter-event time (less standard deviation in

239 Fig. 3d) with lower friction drop. Remark however that, at high particle friction,
240 although the majority of the slip events are smaller, there are also some extreme slip
241 events, which are larger than the events at lower particle friction.

242 When studying the time evolution of the average contact force, we find a similar
243 stick-slip type of behavior (Fig. 6). This observation shows that the contact forces at
244 grain scale control the macroscopic response of the sheared granular gouge. We
245 observe that, by increasing the particle friction, the average contact force increases,
246 however the relative increase of higher particle friction becomes smaller at higher
247 values. The decomposition of contact force into normal and tangential components
248 shows that the contribution of the normal component is dominant irrespective of
249 particle friction value. The relative contribution of normal contact force to the total
250 contact force compared to the contribution of the tangential contact force, however,
251 decreases for higher particle frictions.

252 The Slipping Contact Ratio (SCR) is defined as the ratio between number of
253 contacts at Coulomb frictional limit prone to slip and the total number of contacts.
254 In Fig. 7a, we observe drops in SCR of almost one order of magnitude when
255 changing the particle friction from 0.1 to 0.5 or from 0.5 to 0.9. At the same time,
256 Fig. 7b shows that the average coordination number (coordination number is the
257 number of contacts per particle) decreases with increasing particle friction. The
258 decrease in coordination number for higher particle frictions can be attributed to the
259 larger dilation or higher gouge thickness (see Fig. 2b). We remark that, since the
260 coordination number decreases with increasing particle friction (Fig. 7b), we also

261 check the total number of slipping contacts that is not normalized by the number of
262 contacts and observe similar behavior: the total number of slipping contacts
263 decreases with increasing particle friction.

264 In Figs. 7c and 7d, we show the instantaneous and cumulative particle
265 displacement averaged over all particles, respectively. The cumulative displacement
266 is the total displacement of a particle from the start of each simulation. The
267 instantaneous particle displacement shows larger jumps upon micro- or major slips
268 for higher particle frictions. Moreover, the cumulative particle displacement clearly
269 shows that for a given instant in time (or a given shear strain), particles with higher
270 particle friction have undergone higher total displacements. Our analysis shows that
271 particle displacement is mainly in x-direction (99 %) along the moving boundary that
272 imposes the shear stress (Fig. 1).

273 We also study the evolution of potential and kinetic energies in sheared granular
274 fault gouge (Fig. 8). The elastic strain potential energy is stored within the particle-
275 particle contacts through overlap between particles [Dorostkar & Carmeliet, 2018]
276 and the kinetic energy is due to translation and rotation of particles. The fault gouge
277 with higher particle friction shows more potential energy, however, the increase in
278 average potential energy is not linear with respect to the increase of particle friction.
279 For instance, an increase of particle friction by a factor 9, from 0.1 (Fig. 8c) to 0.9
280 (Fig. 8a), leads to an increase in potential energy only by a factor 3. We also observe
281 that the kinetic energy signal for higher particle frictions shows more fluctuations
282 and bursts i.e. important rearrangements of particles inside the fault gouge. While

283 the potential energy at low particle friction (Fig. 8c) shows a plateau before the major
284 slip events, the energy is always increasing approaching slip events for higher particle
285 frictions (Fig. 8a).

286 We compare in Fig. 9 the evolution of macroscopic friction and gouge thickness
287 for particle frictions with 2 orders of magnitude contrast i.e. 10, 1 and 0.1. The
288 nucleation phase of slip events for higher particle friction (Fig.9a) shows a complex
289 behavior in both macroscopic friction and gouge thickness signals, where a
290 considerable amount of small drops in macroscopic friction occur during the stick
291 phase before an upcoming extreme slip event that has a long recurrence time. The
292 statistical analysis of the size of all slip events based on drop in macroscopic friction
293 (Fig. 10) shows that slip events occur more often for higher particle frictions i.e.
294 larger number of slip events with shorter recurrence interval. More importantly, we
295 observe that, although the higher number of slip events stems mainly from smaller
296 events (see also Fig. 5), there exist also some very large slip (extreme) events (Fig.
297 10).

298 Our analyses show that the velocity (or acceleration) signal of a flagged particle,
299 which is called “pseudo Acoustic Emission (AE)” signal in this study, demonstrates
300 (Fig. 11a) a very similar behavior to acoustic data of the lab (e.g. [Rouet-Leduc et al.,
301 2017]). We will discuss the details of this observation in Section 4. In Fig. 11b, the
302 velocity signal of a flagged particle shows bursts at slip events, which are larger for
303 higher particle frictions. The complementary Cumulative Distribution Function
304 (cCDF) of pseudo acoustic emission bursts (for all emissions without threshold,

305 during both stick and slip phases) in Fig. 12 clearly shows the increase of AE
306 amplitude for higher particle frictions. We remark that the observations in Fig. 11
307 and 12 are not dependent on the flagged particle and are recorded consistently for
308 several chosen particles.

309 We further look at the micromechanics of fault by visualizing the particles on a
310 plane representative of the whole box (here the front x-z plane of the gouge in Fig.
311 1) for particle friction values of 0.1 and 10, at a point during the stick phase (Fig. 13).
312 For panels a and b, we set the maximum of color bar equal to the shear driving plate
313 velocity, 0.06 cm/s. While a gradient is observed for particle friction of 0.1 from top
314 to bottom, where the particles close to the bottom plate have velocity close to 0.06
315 cm/s, the particle velocity field shows a more uniformly distributed profile for a
316 particle friction of 10. We will discuss the implications of this observation in Section
317 4. A comparison between panels c and d shows a larger cumulative displacement for
318 particles with particle friction of 10, where the maximum of color bar is set to the
319 maximum cumulative displacement of particles with particle friction of 0.1, for a
320 better comparison. It is clear from Fig. 13d that a larger portion of the sample has
321 experienced a large displacement. We remind that the main displacement for particles
322 is in x direction, along the driving plate motion. The spatial distributions in panels c
323 and d are consistent with temporal evolution in Fig. 7d. The spatial distributions of
324 coordination number for both particle frictions are rather uniform, while the gouge
325 with lower particle friction shows higher coordination numbers consistent with

326 observations of Fig. 7b. The gouge with higher particle friction shows on average
327 lower numbers, also showing some ‘blue’ spots with very low coordination number.

328 **4- Discussion**

329 We found that the frictional strength of the sheared granular fault gouge during stick-
330 slip dynamics increases with particle friction, but saturates at particle friction values
331 around 0.9-1 (Fig. 3). The nonlinear macroscopic behavior and saturation of friction
332 are attributed to the rotational behavior of the particles showing a gradual transition
333 from sliding to rolling, since high particle friction makes the conditions easier for
334 particles to roll [J. Latham et al., 2005; Azéma et al., 2012; Shojaaee et al., 2012;
335 Göncü & Luding, 2013; Rathbun et al., 2013; Azéma et al., 2017]. Furthermore, we
336 show a linear relation between macroscopic stick-slip friction and gouge thickness
337 (Fig. 4), as is confirmed by previous research [Mead, 1925; C. Marone, 1998; Frye &
338 Marone, 2002; Knuth & Marone, 2007; Makedonska et al., 2011]. From the
339 micromechanical point of view, we only observed a small increase in tangential
340 contact force, while our results show that, the overall shear strength of fault gouge
341 increases at higher particle friction. This means that the increase in shear strength
342 cannot be attributed to the increase of tangential contact force. In other words, the
343 increase in particle friction leads to an increase in normal contact forces enhancing
344 the shear capacity of the system. To explain this observation, we pose the hypothesis
345 that a structural effect is introduced, where at higher particle friction less slipping
346 contacts occur, providing a better support for a contact network to build up higher
347 contact forces. The lower number of slipping contacts (or the higher number of

348 locked contacts) could also explain the larger cumulative particle displacement (Fig.
349 7 and Fig. 13), as rough faults are suggested to require more overall work to shear
350 leading to more deformation in the fault zone [Rathbun et al., 2013]. The hypothesis
351 here is that, the fault gouge with higher particle friction has to undergo more complex
352 states during deformation towards failure, which involves more locked contacts in
353 spite of a lower total number of contacts. This makes the system to experience a
354 wider range of macroscopic frictions and dilations manifested in higher variations of
355 macroscopic friction and gouge thickness (Fig. 3). The system with higher particle
356 friction going through those complex topographical states has to expand more to
357 accommodate the continuous externally applied shear, which leads to a higher
358 dilatation and a lower coordination number. These conditions make a fault gouge
359 with higher particle friction to fail more frequently leading to more fluctuations in
360 both macroscopic friction (slip events) and gouge thickness. A slip event in a sheared
361 granular layer is found to be a phenomenon where the slipping contact ratio increases
362 approaching the failure leading to a major slip event. Therefore, from another point
363 of view, in a system with a higher particle friction and lower slipping contact ratio,
364 the slip event is prevented due to stronger contacts, so that the system will only
365 partially fail leading to a large number of smaller slips. We also remark that, since the
366 main displacement for particles is along the direction of the shear driving plate (x
367 direction in Fig. 1), the higher cumulative particle displacement for a higher particle
368 friction is consistent with a higher number of slip events, where at each slip event
369 there is a displacement (rupture) for the center of mass of the granular fault gouge.

370 Based on our observations, we argue that the effect of gouge particle friction on the
371 mechanics of faults and characteristics of seismic cycles has a similar effect as
372 increasing the fault roughness in absence of fault gouge. The recent numerical model
373 using 2-D plane strain calculations by Tal and Hager [2018] showed that, as the fault
374 surface roughness amplitude increases, the load in the fault is released by more slip
375 events but with lower average stress drops [Tal & Hager, 2018]. Similarly, we observe
376 in Fig. 4 that increase of particle friction leads to more slip events but with lower
377 friction drops. Furthermore, using numerical simulations, Tal et al. [2018] observed
378 a more complex behavior for faults with higher roughness, where the complexities
379 in the nucleation process are reflected as irregular fluctuations in the moment rate
380 for rougher faults [Tal et al., 2018]. We observe a similar complexity in 3D DEM,
381 where the nucleation (stick) phase of slip events contains many fluctuations i.e.
382 smaller slip events. However, using the advantage of DEM and employing the
383 periodic boundary conditions we can shear the fault gouge during long time
384 collecting information of hundreds of slip events. Using a statistical analysis, we then
385 show that a fault gouge with high enough particle friction and large enough shear
386 displacement shows, besides the large amount of small events, some extreme slip
387 events with long recurrence time compared to faults with lower particle friction.

388 Acoustic emission is suggested to originate from groaning, creaking, and
389 chattering of continuous grain motions and breakage of force chains within the fault
390 gouge during laboratory stick-slip experiments [McLaskey & Glaser, 2011; T. H. W.
391 Goebel et al., 2012; T. H. W. Goebel, Candela, et al., 2014; Thomas H.W. Goebel et

392 al., 2017; Rouet-Leduc et al., 2017; Rivière et al., 2018]. We use the velocity signal of
393 flagged particles as pseudo-acoustic emission signals, however, since the velocity
394 profile of a flagged particle contains both motion from arriving waves owing to
395 rearrangements of other particles and the motion of the particle itself, we call it
396 “pseudo AE” signal. In other words, although our AE signal is derived from the
397 motion of a single flagged particle compared to the lab signal where the AE is usually
398 recorded with a device outside of the fault gouge; we find that the pseudo AE signal
399 contains sufficient information for the purpose of this study, where we compare the
400 velocity signal from the same particles for simulations with different particle frictions.
401 As the experimental work by Kwiatek et al. [2014] showed that the observed changes
402 in AE characteristics are clearly correlated to the fault topography and roughness
403 [Kwiatek et al., 2014], Goebel et al. [2017] discussed that faults with rougher surface
404 show a more spatially distributed AE activity and a higher b-value i.e. the measure
405 for the relative abundance of the strong to the weak earthquakes based on the
406 Gutenberg–Richter law [Thomas H.W. Goebel et al., 2017]. Our DEM simulations
407 show dependency of AE on fault particle friction. We observe higher temporal (Fig.
408 11 and 12) and more spatially distributed AE (Fig. 13) for fault with higher particle
409 friction. The higher AE in fault gouge with higher particle friction is consistent with
410 more frequent kinetic energy releases (Fig. 8), since AE is believed to originate from
411 the rearrangement of particles. We remind that, although our discussion on AE is
412 based on velocity tracking of single flagged particles, the temporal evolution and
413 higher moments of those signals show very similar behavior to laboratory AE (e.g.

414 [Rouet-Leduc et al., 2017; Rivière et al., 2018]), as we are using them for machine
415 learning analyses in our ongoing research.

416 Overall, our observations on the effect of particle friction using 3-D DEM model
417 of a granular fault gouge show similarities to other numerical and experimental works
418 on the effect of fault surface roughness, where the presence of granular gouge has
419 been disregarded. In order to better document this observation, in Fig. 14 we show
420 schematically a fictitious shear plane for gouges with low and high particle frictions.
421 For a rough fault with a granular gouge at the core, where the fault surface is totally
422 engaged with the gouge and deformation (slip) takes place inside the gouge zone, our
423 macro- and micro-scale observations suggest high particle friction originates a more
424 complex topography and shear pattern for slip, which is similar to a surface with high
425 roughness amplitude. This rougher shear pattern stems from stronger granular
426 structure at higher particle friction, showing more locked contacts and the larger
427 deformation that the fault gouge has to go through approaching failure. The DEM
428 model in this work is a vast simplification of real faults in nature but yet expands our
429 understanding of micro-scale fault frictional processes and provides a means to study
430 and measure quantities that are not feasible to measure in the lab and in the field,
431 showing how numerical models can boost our understanding from physical
432 processes that dictate frictional strength of a fault damage zone.

433

434

435

436 5- Conclusions

437 We model stick-slip dynamics of a granular fault gouge by 3-D discrete element
438 simulations and study the influence of particle friction on the characteristics of
439 seismic cycles. The major findings of this study can be summarized as follows:

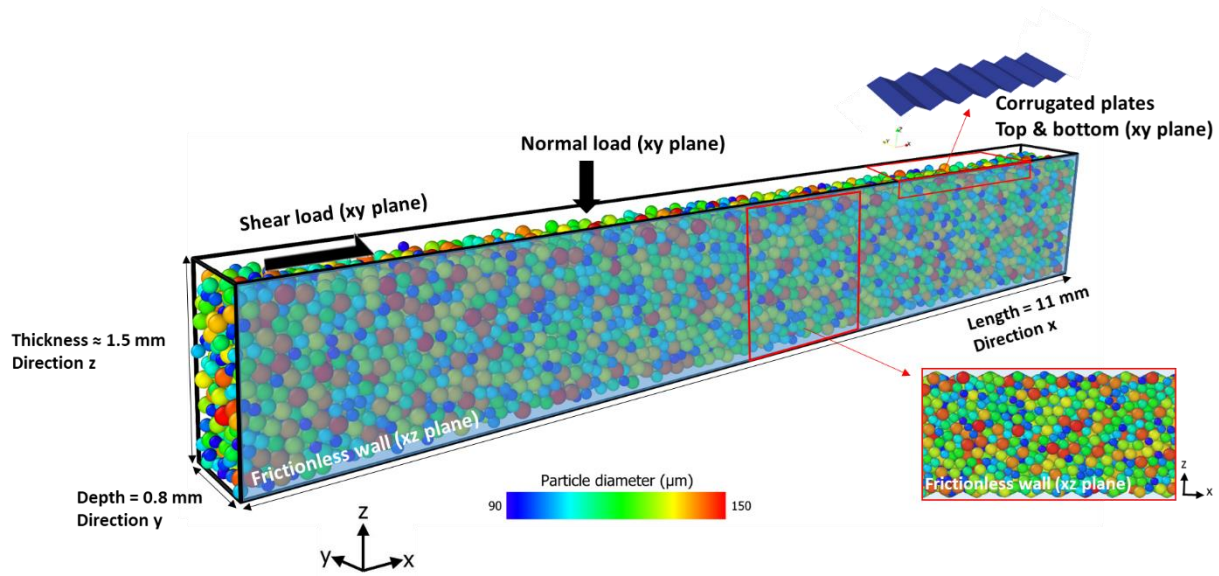
- 440 - The fault gouge frictional strength, dilation and their variations nonlinearly
441 increase with particle friction, saturating at high particle friction.
- 442 - The average slip events' recurrence time and its variations decrease with
443 particle friction. A fault gouge with higher particle friction shows a more
444 complex nucleation (stick) phase, characterized by many smaller slip events as
445 manifested by the more frequent energy release in the kinetic energy signal.
- 446 - Our statistical analyses on a large number of slip events obtained by shearing
447 the fault gouge to a large shear strain suggest that a fault gouge with higher
448 particle friction shows a higher number of slip events mainly consisting in
449 small slip events. However, there are also some extreme slip events larger than
450 the extreme events of fault gouge with lower particle friction.
- 451 - The fault gouge with higher particle friction shows higher stored potential
452 energy and stronger particle-particle contacts leading to a structure that needs
453 more work to deform. Therefore, for a given shear strain, the particles in a
454 gouge with higher particle friction experience more deformation.

- 455 - The pseudo acoustic emission analysis, based on monitoring the velocity of
456 particles, shows higher temporal emissions for fault gouge with higher particle
457 friction.
- 458 - Our observations suggest that, in a rough fault with granular gouge where the
459 rough fault surface is engaged with the gouge zone, the effect of gouge particle
460 friction on the characteristics of seismic cycles is similar to the effect of fault
461 surface roughness in absence of fault gouge.

462 **6- Acknowledgement**

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464 can be obtained by contacting the corresponding author at domid@ethz.ch.

7- Figures

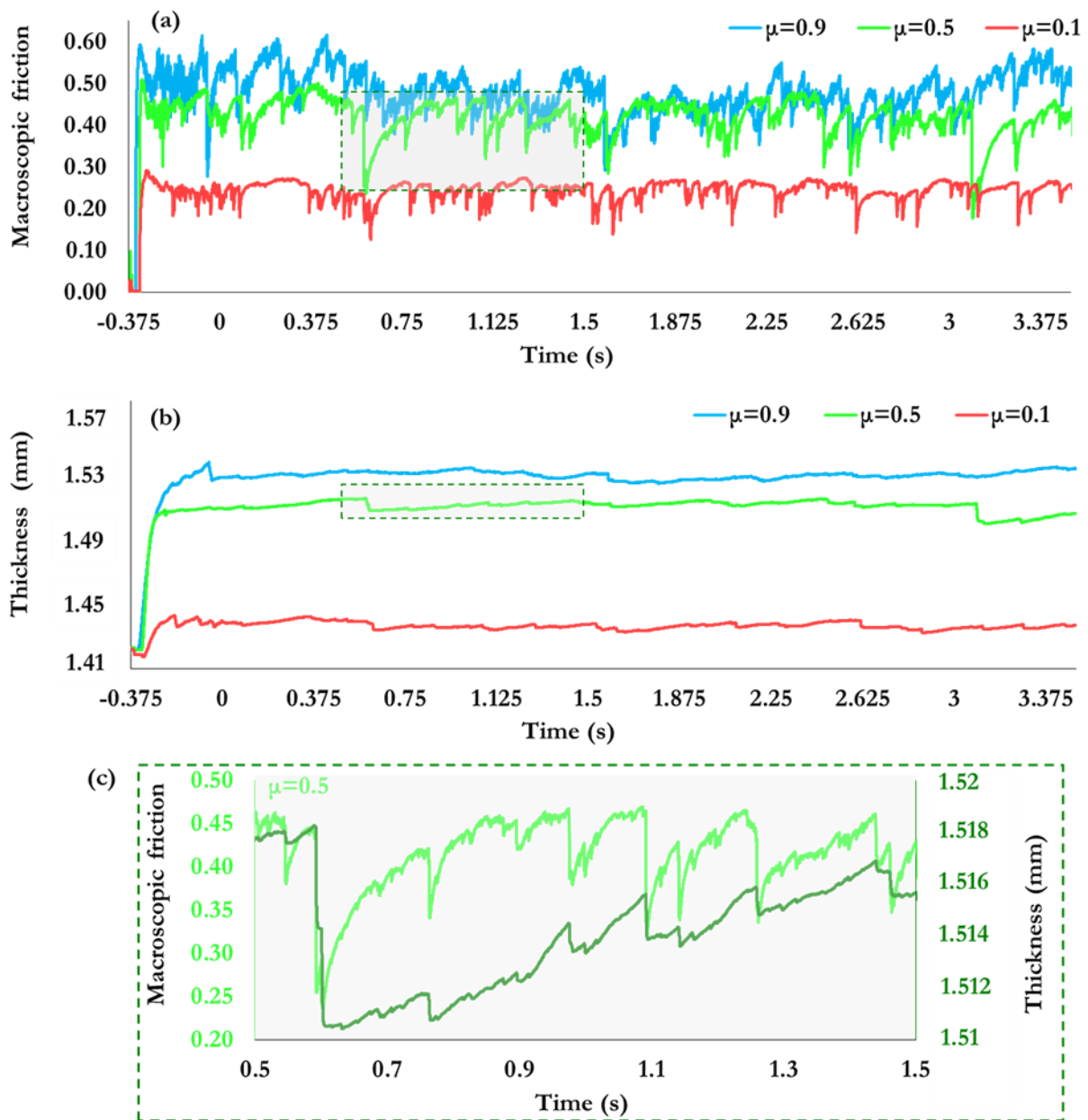


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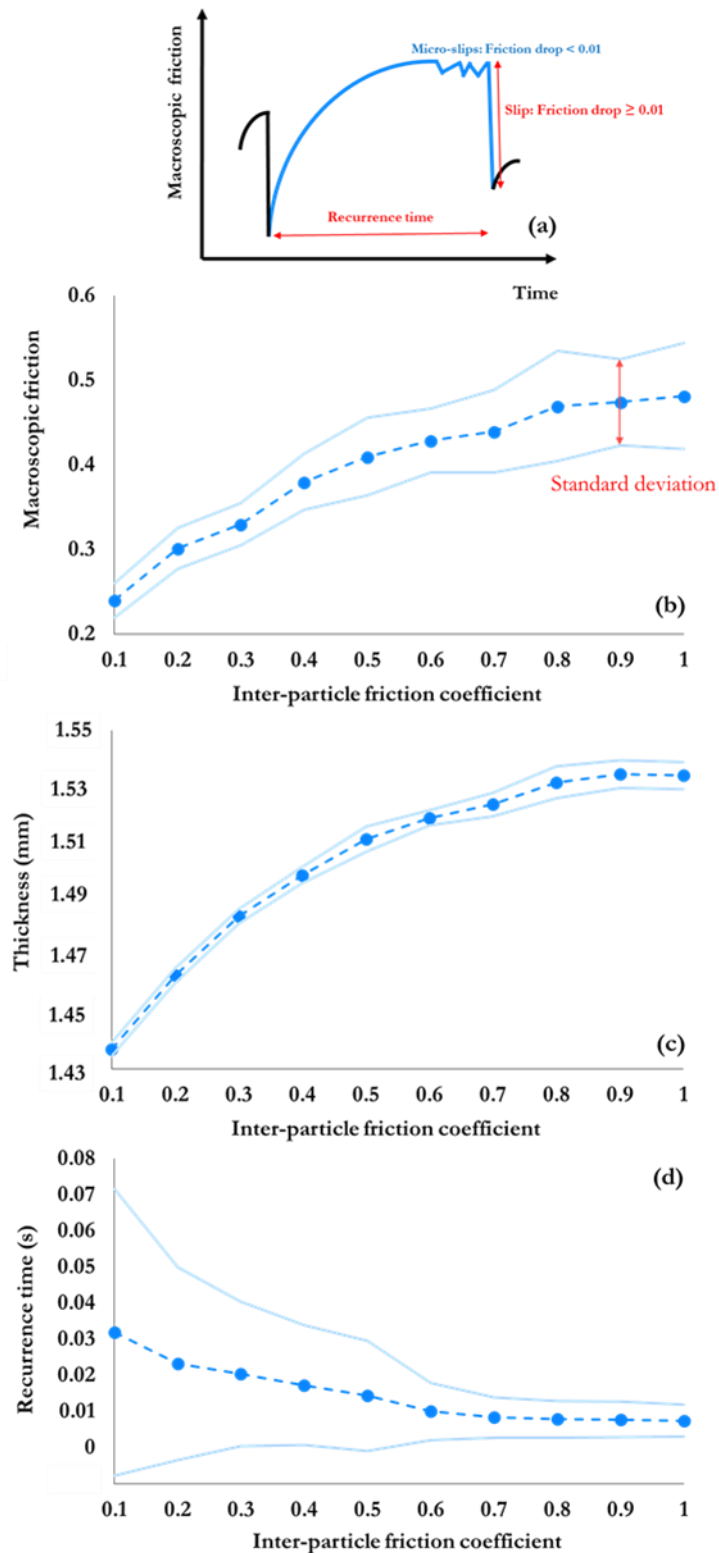
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468

469 **Fig. 1:** Three dimensional granular fault gouge with 8000 particles with poly-disperse
 470 diameter distribution of 90-150 micrometer. The fault gouge is confined in z
 471 direction and sheared in x direction, with periodic boundary conditions at the end
 472 on y-z planes. The shear load is applied along the x-y plane. Two corrugated plates
 473 are used on top and bottom x-y planes of the gouge to simulate a rough fault surface.
 474 The x-z planes are frictionless walls. (Image produced with the open source
 475 visualization tool (OVITO) [Stukowski, 2010]).



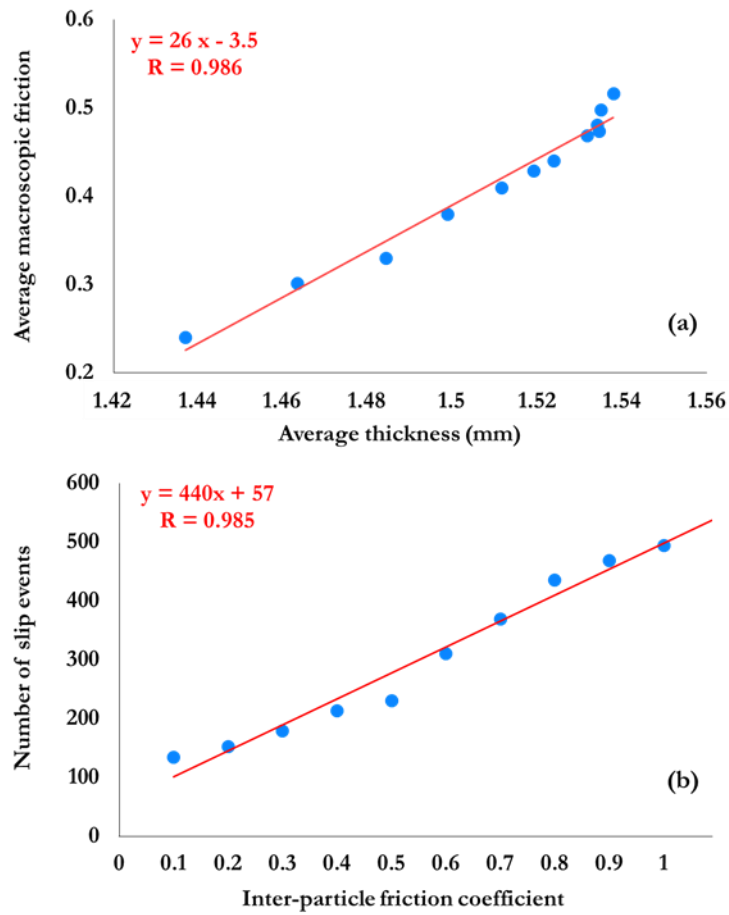
476
 477 **Fig. 2:** Time series of (a) macroscopic friction and (b) fault gouge thickness for three
 478 different particle friction values. The greenish shaded areas in a and b are shown in
 479 (c) with a higher resolution for $\mu = 0.5$. Please note that the thickness is shown with
 480 a secondary axis in c. Time zero denotes the start of stick-slip dynamics after the
 481 initial stable sliding.



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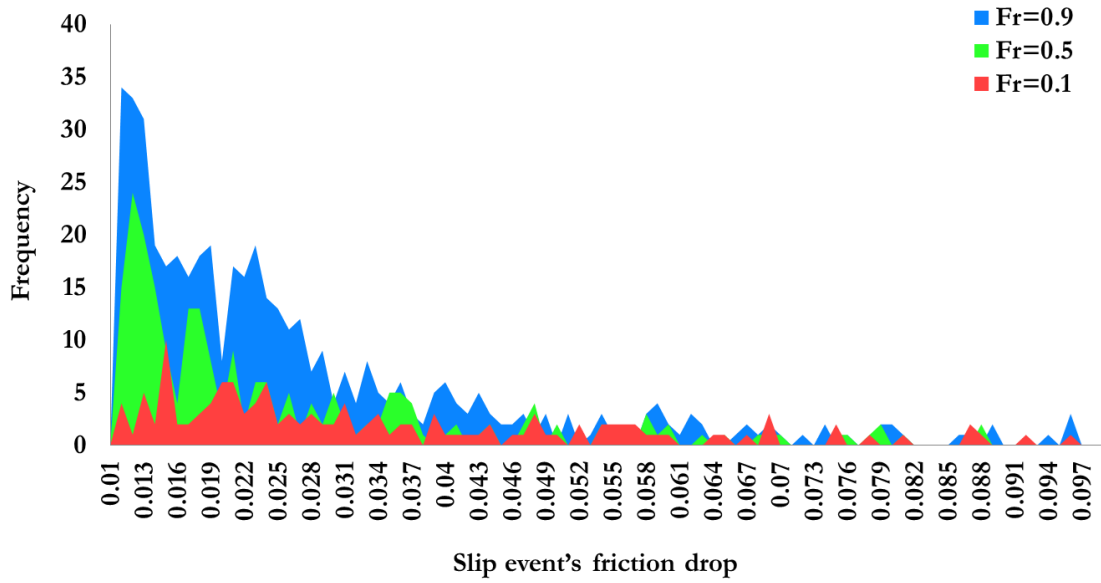
483

484 **Fig. 3:** (a) Schematic showing the definition of slip event, micro-slips and recurrence
 485 interval. Average (b) macroscopic friction, (c) fault gouge thickness and (d) slip
 486 recurrence time as a function of the particle friction. The light blue lines in b-d show
 487 the uncertainty limits based on the standard deviation of data.



488

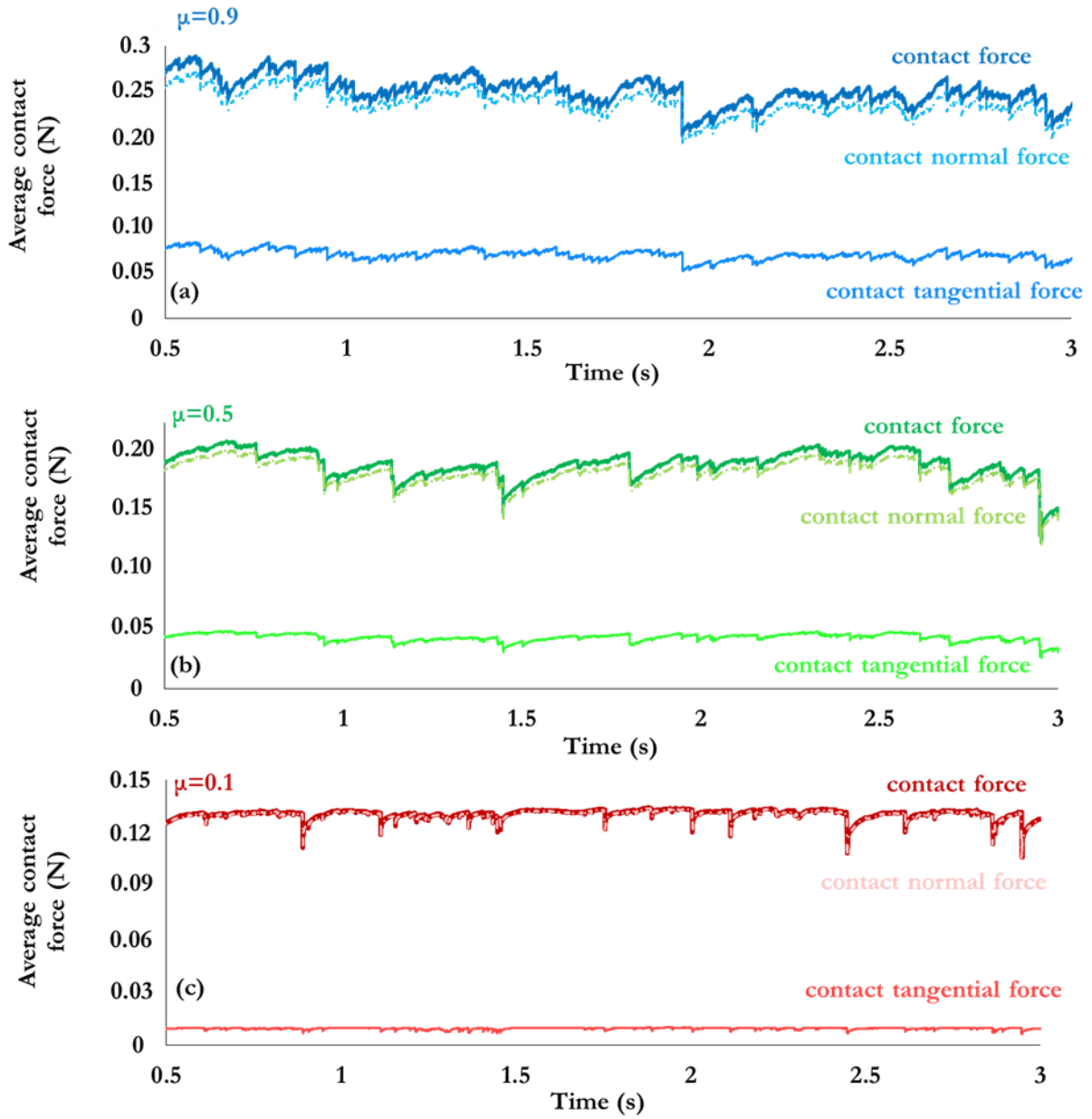
489 **Fig. 4:** (a) Average macroscopic friction versus average thickness for stick-slip
 490 dynamics with different particle friction values. (b) Number of slip events versus
 491 particle friction.



492

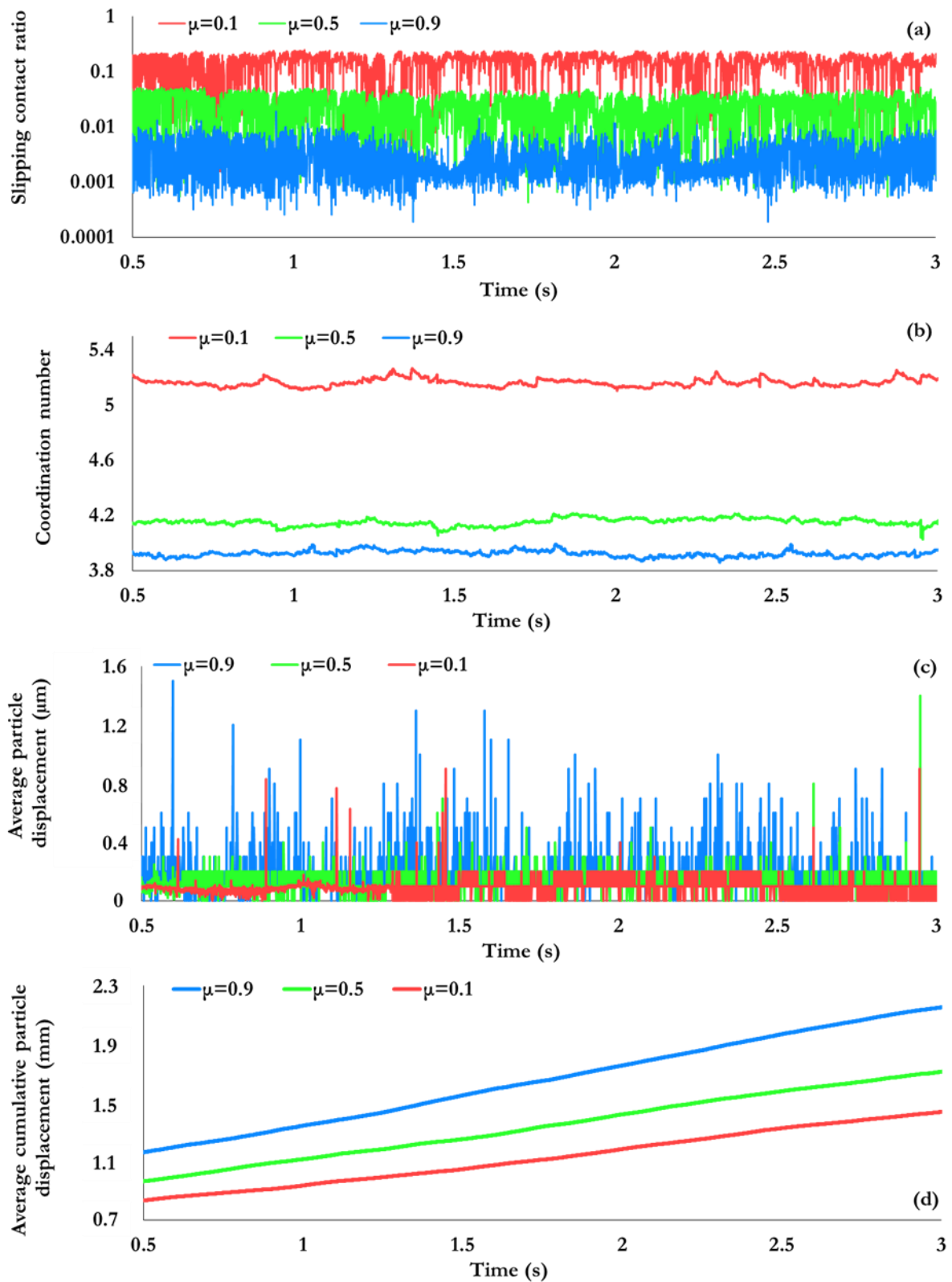
493

494 **Fig. 5:** Histogram of slip event's friction drop for three different particle friction
 495 values. The maximum slip event's friction drop in this histogram is limited to 0.1 to
 496 highlight and better show the frequency of slip events with small friction drop.



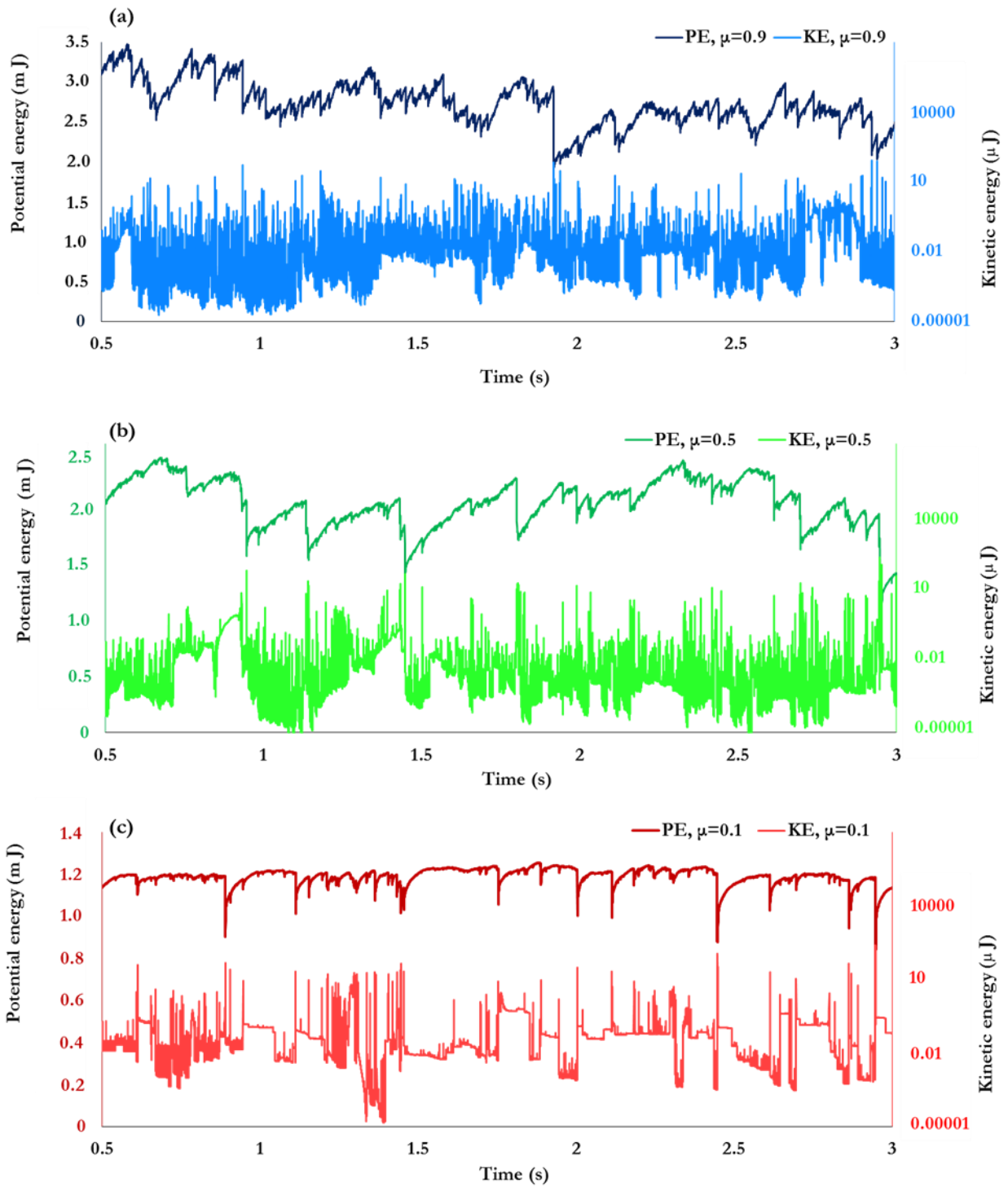
497

498 **Fig. 6:** (a-c) Average contact force for particle friction values of 0.1, 0.5 and 0.9,
 499 respectively. In each panel, the components of contact force (normal contact force
 500 and tangential contact force) are separately shown.



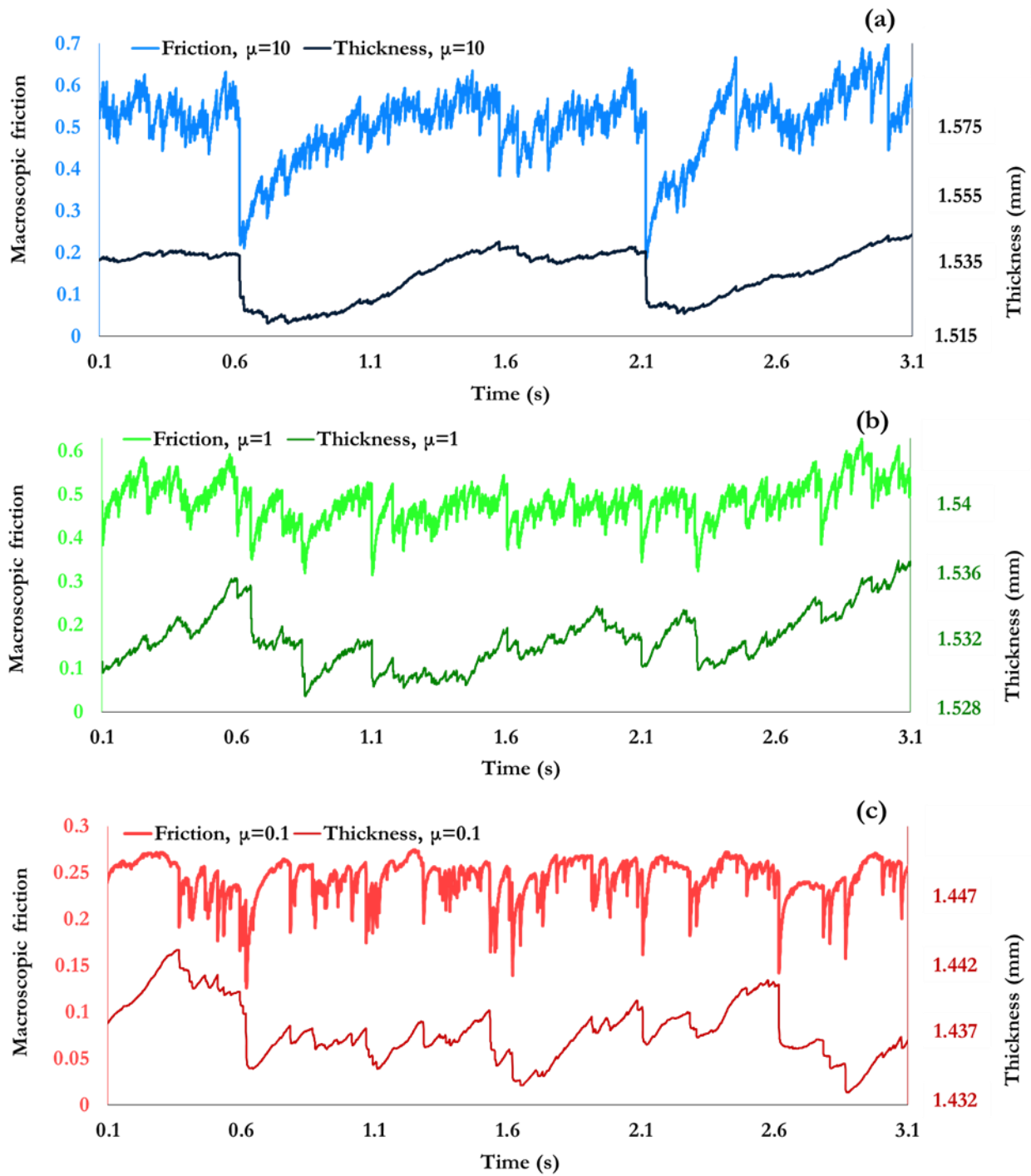
501

502 **Fig. 7:** Time series of (a) slipping contact ratio, (b) coordination number, (c) average
 503 particle displacement and (d) average cumulative particle displacement for particle
 504 friction values of 0.1, 0.5 and 0.9.



505

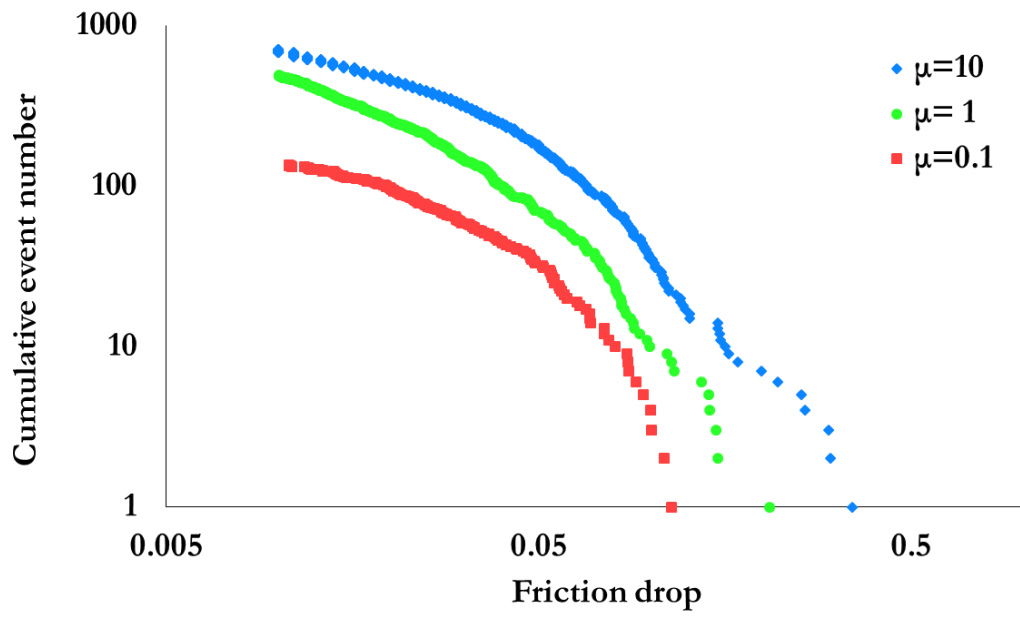
506 **Fig. 8:** (a-c) Evolution of potential energy (primary axis, left) and kinetic energy
 507 (secondary logarithmic axis, right) for particle friction values of 0.9, 0.5 and 0.1,
 508 respectively.



509

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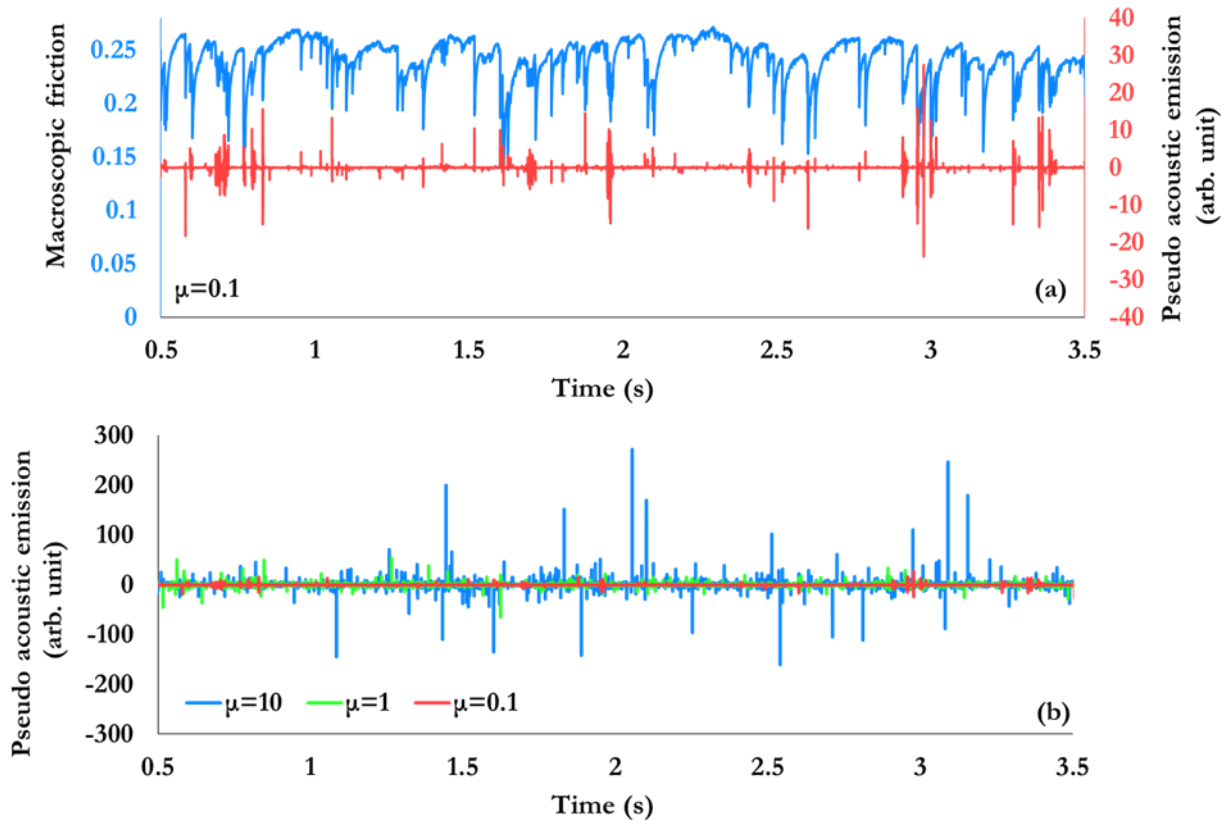
511 **Fig. 9:** (a-c) Time series of macroscopic friction (primary axis, left) and gouge
 512 thickness (secondary axis, right) for particle friction values of 10, 1, 0.1 respectively.



513

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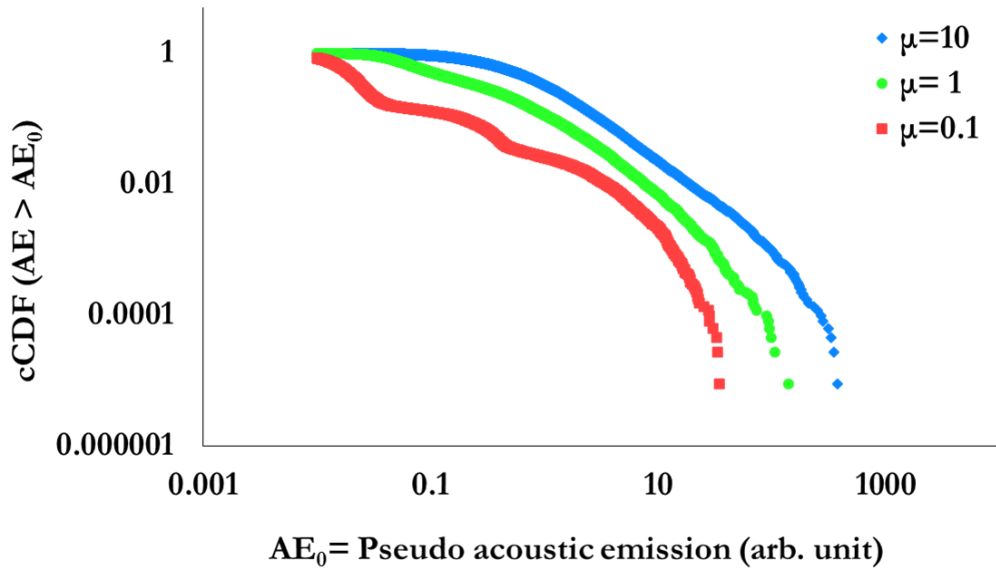
515 **Fig. 10:** Cumulative number of slip events versus macroscopic friction drop for
 516 different particle frictions of 10, 1 and 0.1.



517

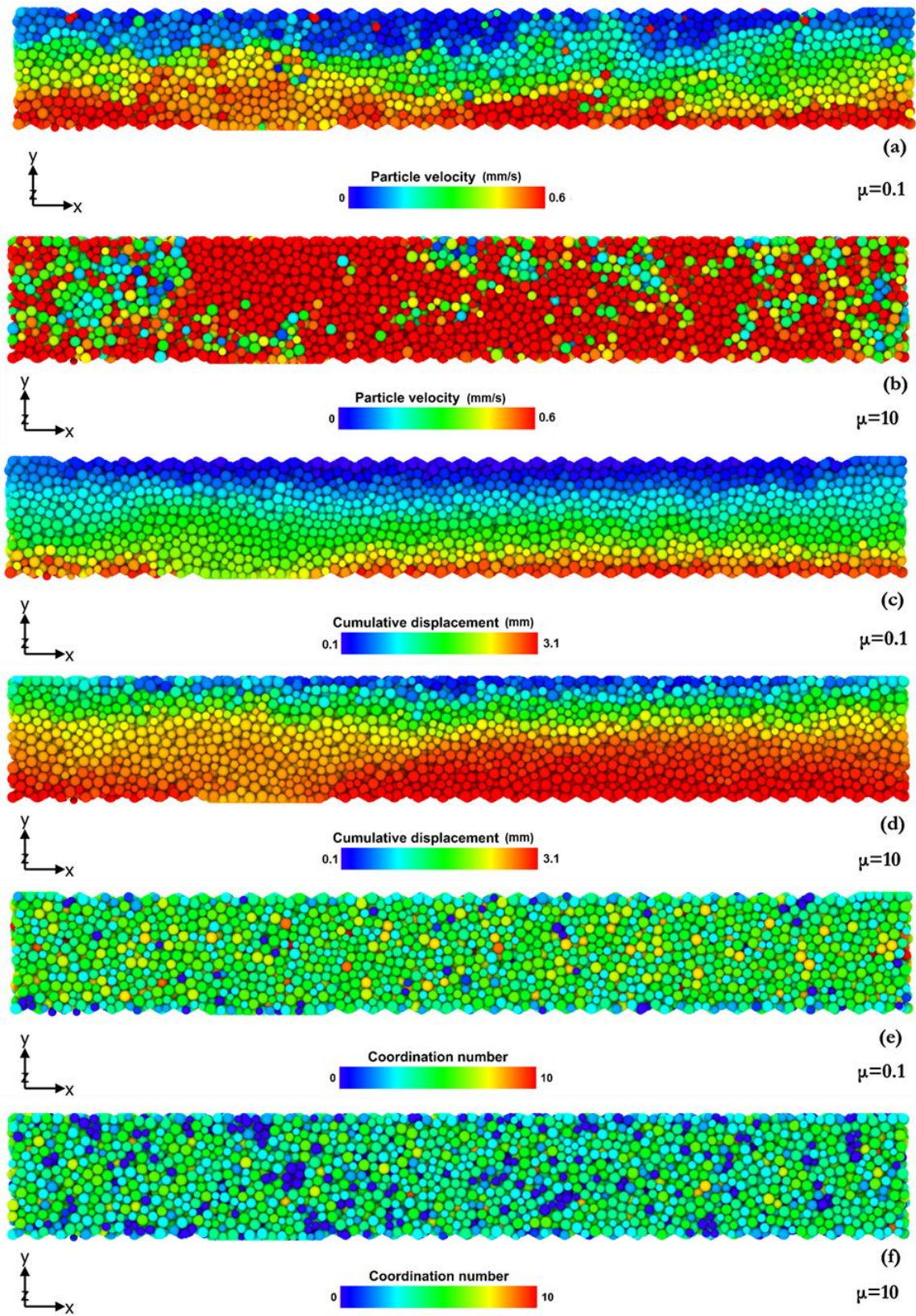
518

519 **Fig. 11:** (a) Macroscopic friction (primary axis, left) and pseudo acoustic emission
 520 (secondary axis, right) for particle friction values of 10, 1 and 0.1. (b) Time series of
 521 pseudo acoustic emission for particle friction values of 10, 1 and 0.1.



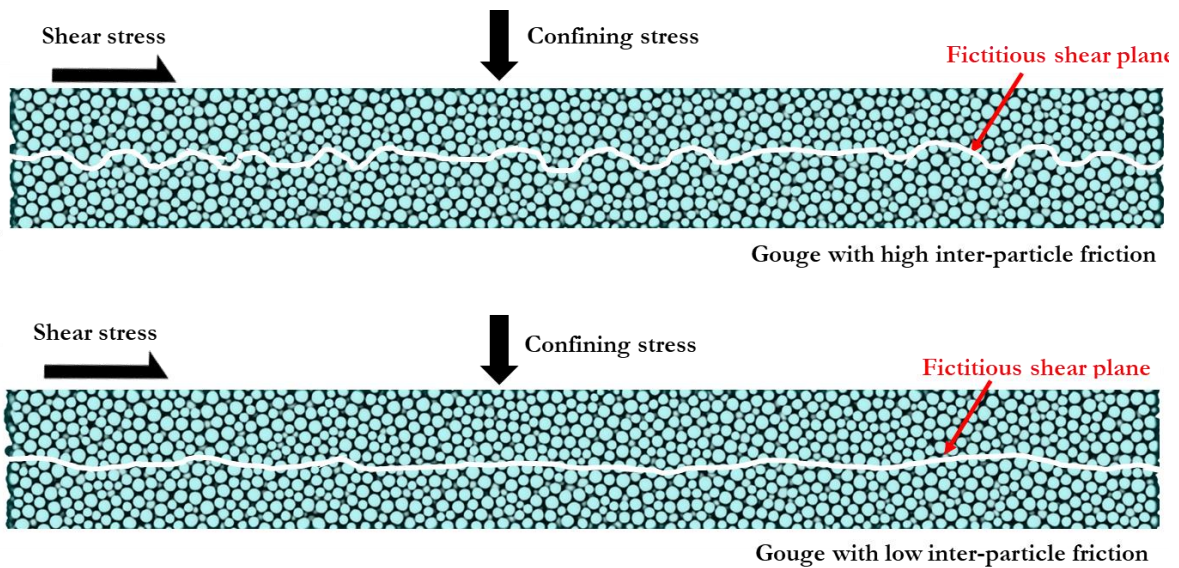
522

523 **Fig. 12:** Complementary Cumulative Distribution Function of pseudo acoustic
 524 emission for different particle friction values of 10, 1 and 0.1.



525

526 **Fig. 13:** (a, c, e) Spatial distribution of particle velocity, cumulative displacement and
 527 coordination number for particle friction values of 0.1 and (b, d, f) for particle
 528 friction of 10, respectively.



529

530 **Fig. 14:** Schematics of a fictitious shear plane at the core of fault gouge with low and
 531 high particle friction. The fictitious shear plane that is shown here with a curved line
 532 is a vast simplification of a complex 3D failure surface or an irregular failure pattern.
 533 The schematics delivers the hypothesis that high particle friction originates a more
 534 complex shear pattern that affects the mechanics of fault and characteristics of
 535 seismic cycles similar to the effect of fault surface roughness in absence of granular
 536 gouge.

8- References

- 537
538
539 Agnolin, I., & Roux, J. N. (2007). Internal states of model isotropic granular packings.
540 I. Assembling process, geometry, and contact networks. *Physical Review E*,
541 76(6). doi:10.1103/PhysRevE.76.061302
- 542 Angheluta, L., Candela, T., Mathiesen, J., & Renard, F. (2011). Effect of Surface
543 Morphology on the Dissipation During Shear and Slip Along a Rock–Rock
544 Interface that Contains a Visco-elastic Core. *Pure and Applied Geophysics*,
545 168(12), 2335-2344. doi:10.1007/s00024-011-0272-8
- 546 Anthony, J. L., & Marone, C. (2005). Influence of particle characteristics on granular
547 friction. *Journal of Geophysical Research-Solid Earth*, 110(B8).
548 doi:10.1029/2004jb003399
- 549 Azéma, E., Estrada, N., Preechawuttipong, I., Delenne, J.-Y., & Radjai, F. (2017).
550 Systematic description of the effect of particle shape on the strength
551 properties of granular media. *EPJ Web Conf.*, 140, 06026.
- 552 Azéma, E., Estrada, N., & Radjaï, F. (2012). Nonlinear effects of particle shape
553 angularity in sheared granular media. *Physical Review E*, 86(4), 041301.
554 doi:10.1103/PhysRevE.86.041301
- 555 Bistacchi, A., Griffith, W. A., Smith, S. A. F., Di Toro, G., Jones, R., & Nielsen, S.
556 (2011). Fault Roughness at Seismogenic Depths from LIDAR and
557 Photogrammetric Analysis. *Pure and Applied Geophysics*, 168(12), 2345-2363.
558 doi:10.1007/s00024-011-0301-7
- 559 Bouchon, M., Karabulut, H., Bouin, M.-P., Schmittbuhl, J., Vallée, M., Archuleta, R.,
560 . . . Marsan, D. (2010). Faulting characteristics of supershear earthquakes.
561 *Tectonophysics*, 493(3), 244-253.
562 doi:<https://doi.org/10.1016/j.tecto.2010.06.011>
- 563 Brantut, N., Schubnel, A., Rouzaud, J. N., Brunet, F., & Shimamoto, T. (2008). High-
564 velocity frictional properties of a clay-bearing fault gouge and implications for
565 earthquake mechanics. *Journal of Geophysical Research: Solid Earth*, 113(B10).
566 doi:10.1029/2007JB005551
- 567 Brodsky, E. E., Gilchrist, J. J., Sagy, A., & Collettini, C. (2011). Faults smooth
568 gradually as a function of slip. *Earth and Planetary Science Letters*, 302(1), 185-
569 193. doi:<https://doi.org/10.1016/j.epsl.2010.12.010>
- 570 Brodsky, E. E., Kirkpatrick, J. D., & Candela, T. (2016). Constraints from fault
571 roughness on the scale-dependent strength of rocks. *Geology*, 44(1), 19-22.

- 572 Bruhat, L., Fang, Z., & Dunham, E. M. (2016). Rupture complexity and the
573 supershear transition on rough faults. *Journal of Geophysical Research: Solid Earth*,
574 121(1), 210-224. doi:doi:10.1002/2015JB012512
- 575 Candela, T., Renard, F., Bouchon, M., Brouste, A., Marsan, D., Schmittbuhl, J., &
576 Voisin, C. (2009). Characterization of Fault Roughness at Various Scales:
577 Implications of Three-Dimensional High Resolution Topography
578 Measurements. *Pure and Applied Geophysics*, 166(10), 1817-1851.
579 doi:10.1007/s00024-009-0521-2
- 580 Candela, T., Renard, F., Bouchon, M., Schmittbuhl, J., & Brodsky, E. E. (2011).
581 Stress drop during earthquakes: effect of fault roughness scaling. *Bulletin of the*
582 *Seismological Society of America*, 101(5), 2369-2387.
- 583 Candela, T., Renard, F., Klinger, Y., Mair, K., Schmittbuhl, J., & Brodsky, E. E.
584 (2012). Roughness of fault surfaces over nine decades of length scales. *Journal*
585 *of Geophysical Research: Solid Earth*, 117(B8). doi:doi:10.1029/2011JB009041
- 586 Candela, T., Renard, F., Schmittbuhl, J., Bouchon, M., & Brodsky, E. E. (2011). Fault
587 slip distribution and fault roughness. *Geophysical Journal International*, 187(2),
588 959-968. doi:10.1111/j.1365-246X.2011.05189.x
- 589 Cashman, S., & Cashman, K. (2000). Cataclasis and deformation-band formation in
590 unconsolidated marine terrace sand, Humboldt County, California. *Geology*,
591 28(2), 111-114. doi:10.1130/0091-7613(2000)28<111:CADFIU>2.0.CO;2
- 592 Cashman, S. M., Baldwin, J. N., Cashman, K. V., Swanson, K., & Crawford, R.
593 (2007). Microstructures developed by coseismic and aseismic faulting in near-
594 surface sediments, San Andreas fault, California. *Geology*, 35(7), 611-614.
- 595 Chester, F. M., & Chester, J. S. (1998). Ultracataclasite structure and friction
596 processes of the Punchbowl fault, San Andreas system, California.
597 *Tectonophysics*, 295(1), 199-221. doi:[https://doi.org/10.1016/S0040-](https://doi.org/10.1016/S0040-1951(98)00121-8)
598 [1951\(98\)00121-8](https://doi.org/10.1016/S0040-1951(98)00121-8)
- 599 Chester, F. M., & Chester, J. S. (2000). Stress and deformation along wavy frictional
600 faults. *Journal of Geophysical Research: Solid Earth*, 105(B10), 23421-23430.
601 doi:doi:10.1029/2000JB900241
- 602 Chester, F. M., Friedman, M., & Logan, J. M. (1985). Foliated cataclasites.
603 *Tectonophysics*, 111(1), 139-146. doi:[https://doi.org/10.1016/0040-](https://doi.org/10.1016/0040-1951(85)90071-X)
604 [1951\(85\)90071-X](https://doi.org/10.1016/0040-1951(85)90071-X)
- 605 Chester, F. M., & Logan, J. M. (1986). Implications for mechanical properties of
606 brittle faults from observations of the Punchbowl fault zone, California. *pure*
607 *and applied geophysics*, 124(1), 79-106. doi:10.1007/bf00875720

- 608 Di Renzo, A., & Di Maio, F. P. (2004). Comparison of contact-force models for the
609 simulation of collisions in DEM-based granular flow codes. *Chemical*
610 *Engineering Science*, 59(3), 525-541. doi:10.1016/j.ces.2003.09.037
- 611 Dieterich, J. H., & Kilgore, B. D. (1994). Direct observation of frictional contacts:
612 New insights for state-dependent properties. *pure and applied geophysics*, 143(1),
613 283-302. doi:10.1007/bf00874332
- 614 Dieterich, J. H., & Smith, D. E. (2009). Nonplanar Faults: Mechanics of Slip and
615 Off-fault Damage. *Pure and Applied Geophysics*, 166(10), 1799-1815.
616 doi:10.1007/s00024-009-0517-y
- 617 Dorostkar, O. (2018). *Stick-slip dynamics in dry and fluid saturated granular fault gouge*
618 *investigated by numerical simulations*. (PhD dissertaion), PhD dissertaion, ETH
619 Zurich.,
- 620 Dorostkar, O., & Carmeliet, J. (2018). Potential Energy as Metric for Understanding
621 Stick–Slip Dynamics in Sheared Granular Fault Gouge: A Coupled CFD–
622 DEM Study. *Rock Mechanics and Rock Engineering*. doi:10.1007/s00603-018-
623 1457-6
- 624 Dorostkar, O., Guyer, R. A., Johnson, P. A., Marone, C., & Carmeliet, J. (2017a). On
625 the micromechanics of slip events in sheared, fluid saturated fault gouge.
626 *Geophysical Research Letters*, 6101–6108. doi:10.1002/2017GL073768
- 627 Dorostkar, O., Guyer, R. A., Johnson, P. A., Marone, C., & Carmeliet, J. (2017b). On
628 the role of fluids in stick-slip dynamics of saturated granular fault gouge using
629 a coupled computational fluid dynamics-discrete element approach. *Journal of*
630 *Geophysical Research: Solid Earth*. doi:10.1002/2017JB014099
- 631 Dorostkar, O., Guyer, R. A., Johnson, P. A., Marone, C., & Carmeliet, J. (2018).
632 Cohesion-induced stabilization in stick-slip dynamics of weakly wet, sheared
633 granular fault gouge. *Journal of Geophysical Research: Solid Earth*.
634 doi:10.1002/2017JB015171
- 635 Dorostkar, O., Johnson, P., Guyer, R., Marone, C., & Carmeliet, J. (2017). Do Fluids
636 Modify the Stick-Slip Behavior of Sheared Granular Media? *Poromechanics VI :*
637 *Proceedings of the Sixth Biot Conference on Poromechanics, 2017.* , 158-163.
638 doi:10.1061/9780784480779.019
- 639 Dunham, E. M., Belanger, D., Cong, L., & Kozdon, J. E. (2011). Earthquake ruptures
640 with strongly rate-weakening friction and off-fault plasticity, Part 2:
641 Nonplanar faults. *Bulletin of the Seismological Society of America*, 101(5), 2308-2322.
- 642 Engelder, J. T. (1974). Cataclasis and the Generation of Fault Gouge. *GSA Bulletin*,
643 85(10), 1515-1522. doi:10.1130/0016-
644 7606(1974)85<1515:CATGOF>2.0.CO;2

- 645 Faulkner, D. R., Lewis, A. C., & Rutter, E. H. (2003). On the internal structure and
646 mechanics of large strike-slip fault zones: field observations of the Carboneras
647 fault in southeastern Spain. *Tectonophysics*, 367(3), 235-251.
648 doi:[https://doi.org/10.1016/S0040-1951\(03\)00134-3](https://doi.org/10.1016/S0040-1951(03)00134-3)
- 649 Fournier, T., & Morgan, J. (2012). Insights to slip behavior on rough faults using
650 discrete element modeling. *Geophysical Research Letters*, 39(12).
651 doi:10.1029/2012GL051899
- 652 Frye, K. M., & Marone, C. (2002). Effect of humidity on granular friction at room
653 temperature. *Journal of Geophysical Research-Solid Earth*, 107(B11).
654 doi:10.1029/2001jb000654
- 655 Goebel, T. H. W., Becker, T. W., Sammis, C. G., Dresen, G., & Schorlemmer, D.
656 (2014). Off-fault damage and acoustic emission distributions during the
657 evolution of structurally complex faults over series of stick-slip events.
658 *Geophysical Journal International*, 197(3), 1705-1718. doi:10.1093/gji/ggu074
- 659 Goebel, T. H. W., Becker, T. W., Schorlemmer, D., Stanchits, S., Sammis, C.,
660 Rybacki, E., & Dresen, G. (2012). Identifying fault heterogeneity through
661 mapping spatial anomalies in acoustic emission statistics. *Journal of Geophysical*
662 *Research: Solid Earth*, 117(B3). doi:10.1029/2011JB008763
- 663 Goebel, T. H. W., Candela, T., Sammis, C. G., Becker, T. W., Dresen, G., &
664 Schorlemmer, D. (2014). Seismic event distributions and off-fault damage
665 during frictional sliding of saw-cut surfaces with pre-defined roughness.
666 *Geophysical Journal International*, 196(1), 612-625. doi:10.1093/gji/ggt401
- 667 Goebel, T. H. W., Kwiatek, G., Becker, T. W., Brodsky, E. E., & Dresen, G. (2017).
668 What allows seismic events to grow big?: Insights from b-value and fault
669 roughness analysis in laboratory stick-slip experiments. *Geology*, 45(9), 815-818.
670 doi:10.1130/G39147.1
- 671 Göncü, F., & Luding, S. (2013). Effect of particle friction and polydispersity on the
672 macroscopic stress–strain relations of granular materials. *Acta Geotechnica*, 8(6),
673 629-643. doi:10.1007/s11440-013-0258-z
- 674 Goniva, C., Kloss, C., Deen, N. G., Kuipers, J. A. M., & Pirker, S. (2012). Influence
675 of rolling friction on single spout fluidized bed simulation. *Particuology*, 10(5),
676 582-591. doi:10.1016/j.partic.2012.05.002
- 677 Haines, S., Marone, C., & Saffer, D. (2014). Frictional properties of low-angle normal
678 fault gouges and implications for low-angle normal fault slip. *Earth and*
679 *Planetary Science Letters*, 408, 57-65. doi:10.1016/j.epsl.2014.09.034
- 680 Heermance, R., Shipton, Z. K., & Evans, J. P. (2003). Fault Structure Control on
681 Fault Slip and Ground Motion during the 1999 Rupture of the Chelungpu

- 682 Fault, Taiwan. *Bulletin of the Seismological Society of America*, 93(3), 1034-1050.
683 doi:10.1785/0120010230
- 684 Hertz, H. (1882). Ueber die Berührung fester elastischer Körper. *Journal für die reine*
685 *und angewandte Mathematik*, 92, 156-171.
- 686 J. Latham, S., Abe, S., & R. Mora, P. (2005). *Macroscopic Friction Response of Rotational*
687 *and Non-Rotational Lattice Solid Gouge Models in 2D and 3D*. Paper presented at
688 the Powders and Grains 2005, Stuttgart, Germany.
- 689 Karner, S. L., & Marone, C. (2000). Effects of Loading Rate and Normal Stress on
690 Stress Drop and Stick-Slip Recurrence Interval. *GeoComplexity and the Physics of*
691 *Earthquakes*, 187-198.
- 692 Kloss, C., Goniva, C., Hager, A., Amberger, S., & Pirker, S. (2012). Models,
693 algorithms and validation for opensource DEM and CFD-DEM. *Progress in*
694 *Computational Fluid Dynamics*, 12(2-3), 140-152.
- 695 Knuth, M., & Marone, C. (2007). Friction of sheared granular layers: Role of particle
696 dimensionality, surface roughness, and material properties. *Geochemistry,*
697 *Geophysics, Geosystems*, 8(3). doi:doi:10.1029/2006GC001327
- 698 Kozłowska, M., Brudzinski, M. R., Friberg, P., Skoumal, R. J., Baxter, N. D., &
699 Currie, B. S. (2018). Maturity of nearby faults influences seismic hazard from
700 hydraulic fracturing. *Proceedings of the National Academy of Sciences*, 115(8), E1720-
701 E1729. doi:10.1073/pnas.1715284115
- 702 Kwiatek, G., Goebel, T. H. W., & Dresen, G. (2014). Seismic moment tensor and b
703 value variations over successive seismic cycles in laboratory stick-slip
704 experiments. *Geophysical Research Letters*, 41(16), 5838-5846.
705 doi:doi:10.1002/2014GL060159
- 706 Mair, K., Frye, K. M., & Marone, C. (2002). Influence of grain characteristics on the
707 friction of granular shear zones. *Journal of Geophysical Research-Solid Earth*,
708 107(B10). doi:10.1029/2001jb000516
- 709 Makedonska, N., Sparks, D. W., Aharonov, E., & Goren, L. (2011). Friction versus
710 dilation revisited: Insights from theoretical and numerical models. *Journal of*
711 *Geophysical Research*, 116(B9). doi:10.1029/2010jb008139
- 712 Marone, C. (1998). The effect of loading rate on static friction and the rate of fault
713 healing during the earthquake cycle. *Nature*, 391(6662), 69-72.
- 714 Marone, C. (1998). Laboratory-derived friction laws and their application to seismic
715 faulting. *Annual Review of Earth and Planetary Sciences*, 26, 643-696.
716 doi:10.1146/annurev.earth.26.1.643

- 717 Marone, C., Raleigh, C. B., & Scholz, C. H. (1990). Frictional behavior and
718 constitutive modeling of simulated fault gouge. *Journal of Geophysical Research-
719 Solid Earth and Planets*, 95(B5), 7007-7025. doi:10.1029/JB095iB05p07007
- 720 McLaskey, G. C., & Glaser, S. D. (2011). Micromechanics of asperity rupture during
721 laboratory stick slip experiments. *Geophysical Research Letters*, 38(12).
722 doi:doi:10.1029/2011GL047507
- 723 Mead, W. J. (1925). The geologic role of dilatancy. *The Journal of Geology*, 33(7), 685-
724 698.
- 725 MiDi, G. D. R. (2004). On dense granular flows. *European Physical Journal E*, 14(4),
726 341-365. doi:10.1140/epje/i2003-10153-0
- 727 Ohnaka, M. (2003). A constitutive scaling law and a unified comprehension for
728 frictional slip failure, shear fracture of intact rock, and earthquake rupture.
729 *Journal of Geophysical Research: Solid Earth*, 108(B2).
730 doi:doi:10.1029/2000JB000123
- 731 Ohnaka, M., & Shen, L.-f. (1999). Scaling of the shear rupture process from
732 nucleation to dynamic propagation: Implications of geometric irregularity of
733 the rupturing surfaces. *Journal of Geophysical Research: Solid Earth*, 104(B1), 817-
734 844. doi:doi:10.1029/1998JB900007
- 735 Rathbun, A. P., & Marone, C. (2010). Effect of strain localization on frictional
736 behavior of sheared granular materials. *Journal of Geophysical Research*, 115(B1).
737 doi:10.1029/2009jb006466
- 738 Rathbun, A. P., Renard, F., & Abe, S. (2013). Numerical investigation of the interplay
739 between wall geometry and friction in granular fault gouge. *Journal of Geophysical
740 Research: Solid Earth*, 118(3), 878-896. doi:doi:10.1002/jgrb.50106
- 741 Renard, F., Voisin, C., Marsan, D., & Schmittbuhl, J. (2006). High resolution 3D laser
742 scanner measurements of a strike-slip fault quantify its morphological
743 anisotropy at all scales. *Geophysical Research Letters*, 33(4).
744 doi:doi:10.1029/2005GL025038
- 745 Rijsingen, E., Lallemand, S., Peyret, M., Arcay, D., Heuret, A., Funicello, F., & Corbi,
746 F. (2018). How subduction interface roughness influences the occurrence of
747 large interplate earthquakes. *Geochemistry, Geophysics, Geosystems*.
748 doi:doi:10.1029/2018GC007618
- 749 Rivière, J., Lv, Z., Johnson, P. A., & Marone, C. (2018). Evolution of b-value during
750 the seismic cycle: Insights from laboratory experiments on simulated faults.
751 *Earth and Planetary Science Letters*, 482, 407-413.
752 doi:<https://doi.org/10.1016/j.epsl.2017.11.036>

- 753 Rosenau, M., Corbi, F., & Dominguez, S. (2017). Analogue earthquakes and seismic
754 cycles: experimental modelling across timescales. *Solid Earth*, 8(3), 597-635.
755 doi:10.5194/se-8-597-2017
- 756 Rouet-Leduc, B., Hulbert, C., Lubbers, N., Barros, K., Humphreys, C. J., & Johnson,
757 P. A. (2017). Machine Learning Predicts Laboratory Earthquakes. *Geophysical*
758 *Research Letters*, 44(18), 9276-9282. doi:10.1002/2017GL074677
- 759 Sagy, A., Brodsky, E. E., & Axen, G. J. (2007). Evolution of fault-surface roughness
760 with slip. *Geology*, 35(3), 283-286. doi:10.1130/G23235A.1
- 761 Sheng, Y., Lawrence, C. J., Briscoe, B. J., & Thornton, C. (2004). Numerical studies
762 of uniaxial powder compaction process by 3D DEM. *Engineering Computations*,
763 21(2-4), 304-317. doi:10.1108/02644400410519802
- 764 Shimamoto, T. (1979). *Experimental studies of simulated gouge and their application to studies*
765 *of natural fault zones*. Paper presented at the Conference VIII: Analysis of Actual
766 Fault Zones in Bedrock, Natl. Earthquake Hazards Reduct. Program, Menlo
767 Park, Calif.
- 768 Shojaee, Z., Brendel, L., Török, J., & Wolf, D. E. (2012). Shear flow of dense
769 granular materials near smooth walls. II. Block formation and suppression of
770 slip by rolling friction. *Physical Review E*, 86(1), 011302.
771 doi:10.1103/PhysRevE.86.011302
- 772 Tal, Y., & Hager, B. H. (2018). The Slip Behavior and Source Parameters for
773 Spontaneous Slip Events on Rough Faults Subjected to Slow Tectonic
774 Loading. *Journal of Geophysical Research: Solid Earth*, 123(2), 1810-1823.
775 doi:doi:10.1002/2017JB014737
- 776 Tal, Y., Hager, B. H., & Ampuero, J. P. (2018). The Effects of Fault Roughness on
777 the Earthquake Nucleation Process. *Journal of Geophysical Research: Solid Earth*,
778 123(1), 437-456. doi:doi:10.1002/2017JB014746
- 779 Wang, K., & Bilek, S. L. (2014). Invited review paper: Fault creep caused by
780 subduction of rough seafloor relief. *Tectonophysics*, 610, 1-24.
781 doi:<https://doi.org/10.1016/j.tecto.2013.11.024>
- 782 Zielke, O., Galis, M., & Mai, P. M. (2017). Fault roughness and strength
783 heterogeneity control earthquake size and stress drop. *Geophysical Research*
784 *Letters*, 44(2), 777-783. doi:doi:10.1002/2016GL071700
- 785 Zoback, M., Hickman, S., & Ellsworth, W. (2010). Scientific drilling into the San
786 Andreas fault zone. *Eos, Transactions American Geophysical Union*, 91(22), 197-
787 199.