

1           **GRAIN FRICTION CONTROLS CHARACTERISTICS OF**  
2                   **SEISMIC CYCLE IN GRANULAR FAULT GOUGE;**  
3           **IMPLICATIONS FOR FAULT PARTICLE ROUGHNESS**

4   **Authors:**

5   **I.   Omid Dorostkar<sup>1,2,\*</sup>**

- 6       1. Chair of Building Physics, Department of Mechanical and Process  
7           Engineering, Swiss Federal Institute of Technology Zurich (ETH Zurich),  
8           Stefano-Franscini-Platz 5, CH-8093 Zurich, Switzerland  
9       2. Department of Civil, Environmental and Geomatic Engineering, Swiss  
10           Federal Institute of Technology Zurich (ETH Zurich), Stefano-Franscini-  
11           Platz 5, CH-8093 Zurich, Switzerland

12   \*Corresponding author:

13    Email: [domid@ethz.ch](mailto:domid@ethz.ch)      Phone: +41587654106

14    ID: [orcid.org/0000-0002-7758-4919](https://orcid.org/0000-0002-7758-4919)

15   **II.   Jan Carmeliet<sup>1</sup>**

- 16       1. Chair of Building Physics, Department of Mechanical and Process  
17           Engineering, Swiss Federal Institute of Technology Zurich (ETH Zurich),  
18           Stefano-Franscini-Platz 5, CH-8093 Zurich, Switzerland

19    Email: [cajan@ethz.ch](mailto:cajan@ethz.ch)

25 **Abstract:**

26 In this paper, we study the effect of particle roughness approximated by the inter-  
27 particle friction coefficient on the characteristics of seismic cycles. Our discrete  
28 element simulations show that the stick-slip frictional strength and dilation of the  
29 fault gouge, as well as their variability, nonlinearly increase with the particle  
30 roughness (inter-particle friction coefficient) but at high particle roughness saturate.  
31 By statistical analyses on a large number of slip events, we find that the average  
32 recurrence time and its variability decreases with particle roughness. A rougher fault  
33 shows a more complex nucleation phase and more frequent energy releases in the  
34 kinetic energy signal. We find that the rougher fault on average shows more small  
35 slip events compared to smoother fault but also contains a limited number of  
36 extreme events. The fault gouge with higher particle surface roughness shows higher  
37 stored potential energy and stronger particle-particle contact forces. Our pseudo  
38 acoustic emission analysis, based on the monitoring of the velocity signal of particles,  
39 shows higher temporal and more spatially distributed acoustic emissions for fault  
40 gouge with higher particle surface roughness. Our findings in this study show that  
41 roughness at micro-scale plays an important role in nucleation and rupture process  
42 of earthquakes, similar to conditions where no granular gouge was considered. Our  
43 results in this study are consistent with previous numerical and experimental works  
44 and complete them by focusing on micromechanics of fault damage zone, showing  
45 how numerical models and in particular discrete element simulations can help  
46 enhance our understanding from fault mechanics.

47 **Keywords:** friction, roughness, stick-slip, granular materials, fault gouge, fault  
48 mechanics

## 49        **1- Introduction**

50    The stability and slip behavior of natural faults and their relation with fault properties  
51    and in particular, their surface roughness, have been under debate during last years.  
52    The fault roughness spans from microns to tens of kilometers and can be measured  
53    using different methods. Brodsky et al. [2016] quantified the slip surface roughness  
54    and found that fault surface is rougher at small scales than large ones and that the  
55    scale dependence of roughness implies yielding of asperities at all scales [*Brodsky et*  
56    *al., 2016*]. Using three independent scanner devices, Candela et al. [2012]  
57    characterized the roughness of fault surfaces by a single anisotropic self-affine  
58    description [*Candela et al., 2012*]. The fault surface roughness has been evaluated also  
59    by measurements on exhumed faults [*Renard et al., 2006; Sagy et al., 2007; Candela et*  
60    *al., 2009; Bistacchi et al., 2011; Brodsky et al., 2011*].

61        At field scale, the fault roughness is suggested to control the stress drop and slip  
62    distribution during earthquakes [*Bouchon et al., 2010; Candela et al., 2011a; Candela et*  
63    *al., 2011b*]. The maturity and roughness level of nearby faults are also reported to  
64    influence seismic hazard associated with hydraulic fracturing [*Kozłowska et al., 2018*].  
65    A recent report shows that the seafloor in front of large earthquakes is generally  
66    smoother than in areas where no large earthquakes have occurred [*Rijsingen et al.,*  
67    *2018*]. In their review paper on field scale observation regarding fault creep caused  
68    by subduction of rough seafloor relief, Wang & Bilek [2014] reported lack of  
69    evidence for rough faults to be more strongly locked, while creeping is observed for  
70    both smooth and rough faults [*Wang & Bilek, 2014*].

71 It is barely feasible to study the relation of roughness and seismicity in nature. At lab  
72 scale, experimental studies showed that roughness promotes a larger nucleation  
73 segment [Obnaka & Shen, 1999; Obnaka, 2003]. The study by Goebel et al. [2014]  
74 showed that, as the shearing of an experimental sample advances, there exists a rapid  
75 spatial decay of acoustic emission events during following inter-slip periods, that is  
76 an intimation of decreasing fault zone complexity and fault surface roughness [Goebel  
77 et al., 2014a]. In experiments, the smoother faults are proposed to show super-shear  
78 ruptures as roughness controls the velocity of rupture propagation [Xia et al., 2004;  
79 Schubnel et al., 2011]. The topology and fault structure influence the spatial and  
80 temporal distribution of small and large earthquakes. Using laboratory experiments,  
81 Goebel et al. [2017] showed that more mature, smoother faults produce localized  
82 seismicity, smoother stress fields, and lower b-values (the measure for the relative  
83 abundance of the strong to the weak earthquakes based on the Gutenberg–Richter  
84 law), and that the acoustic emission statistics during stick-slip are strongly influenced  
85 by fault roughness [Goebel et al., 2017]. In their recent review on scaling of fault  
86 roughness and its implications for earthquake mechanics, Renard & Candela [2017]  
87 emphasized on the open question of fault roughness at small scales (below the  
88 millimeter scale) and its possible implications for seismic and aseismic slip [Renard &  
89 Candela, 2017].

90 Numerical simulations have been also used to study the effect of roughness on  
91 mechanics and nucleation of slip events as well as heterogeneity of stress distribution  
92 and dissipation due to fault zone roughness [Chester & Chester, 2000; Dieterich & Smith,

93 2009; *Anghelescu et al., 2011; Dunham et al., 2011; Bruhat et al., 2016; Tal & Hager, 2018;*  
94 *Tal et al., 2018*]. Previous research using 2-D discrete element method (DEM)  
95 showed that the fault roughness can control the slip style from stick-slip to stable  
96 sliding [*Fournier & Morgan, 2012*]. Using 3-D DEM, Rathbun et al. [2013] found that,  
97 in stable sliding regime, roughness and thus coupling of a fault to the gouge zones  
98 influences the number of sliding contacts controlling the critical slip distance [*Marone*  
99 *& Kilgore, 1993; Rathbun et al., 2013*]. Using large-scale numerical simulations, Zielke  
100 et. al [2017] suggested that, besides stress drop and rupture dimension, the  
101 earthquake moment release and its recurrence probability is dependent also on  
102 surface roughness [*Zielke et al., 2017*]. In a new numerical study, Pierre et al. [2018]  
103 showed, without appealing to complex friction rheology, that a simple geometrical  
104 complexity with two overlapping faults give rise to both slow and fast slip events  
105 [*Pierre et al., 2018*].

106       Faults at their core are usually characterized by a granular gouge layer created by  
107 wear, comminution and other frictional processes, as observed in nature [*Engelder,*  
108 *1974; Shimamoto, 1979; Chester et al., 1985; Chester & Logan, 1986; Marone et al., 1990;*  
109 *Chester & Chester, 1998; Cashman & Cashman, 2000; Chester & Chester, 2000; Faulkner*  
110 *et al., 2003; Heermance et al., 2003; Cashman et al., 2007; Zoback et al., 2010*].  
111 Characteristics of slip events in laboratory stick-slip experiments including stress  
112 drop, pre-seismic friction, recurrence time, granular layer compaction etc. have been  
113 studied widely for different loading configurations and particle properties [*Dieterich*  
114 *& Kilgore, 1994; Marone, 1998a; Karner & Marone, 2000; Mair et al., 2002; Brantut et al.,*

115 2008; Rathbun & Marone, 2010; Haines et al., 2014; Rosenau et al., 2017], where the  
116 frictional strength and stability of a sheared granular zone are found sensitive to grain  
117 shape, particle size distribution and their evolution [Mair et al., 2002; Anthony &  
118 Marone, 2005].

119 Despite its dependency on many factors, generally the contact friction coefficient  
120 increases with surface roughness parameters [Biegel et al., 1989; Ivković et al., 2000; Park  
121 & Song, 2009]. In this study, we systematically vary the inter-particle friction  
122 coefficient as a proxy for particle roughness in a granular fault gouge, as the presence  
123 of granular fault gouge is neglected in most of the previous studies on fault  
124 roughness. We focus on stick-slip dynamics and the effect of particle roughness on  
125 characteristics of seismic cycles to answer the following question; “what is the effect  
126 of inter-particle roughness on the slip size distribution and inter-event time of seismic  
127 cycles in a fault with granular fault gouge?”. To this end, we perform 3-D DEM  
128 simulations recording hundreds of slip events for statistical analyses provided by the  
129 advantage of numerical simulations. We also study the evolution of elastic strain  
130 energy (potential energy [Dorostkar, 2018; Dorostkar & Carmeliet, 2018]), kinetic  
131 energy, and the micromechanics of slip events and explain the macro-scale response  
132 of fault with grain-scale metrics (e.g. [Dorostkar et al., 2017a]). In absence of complex  
133 rheological and geometrical granular contact model, we will show that, although our  
134 simplified numerical approach does not contain many complexities present in nature,  
135 it can yet show similar observations to recent numerical and experimental studies on  
136 fault roughness and complete them by statistical analyses using a large quantity of

137 slip events. Our findings in this study will show that roughness at micro-scale level,  
138 approximated by inter-particle friction coefficient, may play a similar role in  
139 mechanics of earthquakes as roughness of fault blocks' surface in absence of fault  
140 gouge.

## 141 **2- Model description**

142 Due to particulate nature of granular fault gouge, we use DEM to model gouge  
143 grains. In DEM, the equations of motion solved for each particle considering the  
144 applied forces are:

$$145 \quad \sum F_p = m \left( \frac{d}{dt} u_p \right), \quad (1)$$

$$146 \quad \sum T_p = I \left( \frac{d}{dt} \omega_p \right), \quad (2)$$

147 where  $m$ ,  $I$ ,  $u_p$  and  $\omega_p$  are the mass, the moment of inertia and the translational and  
148 angular velocity of particle, respectively. In Eq. (1) and (2),  $F_p$  and  $T_p$  are the forces  
149 and torques acting on particle  $i$  owing to particle-particle contacts. In soft sphere  
150 DEM, the particle-particle contact allows an overlap between them and the contact  
151 law is described by a combination of different rheological elements (spring, dashpot,  
152 slider etc.). These elements are at play when particles are in contact. Upon contact  
153 loss, there are no contact forces acting between the particles. The process of forming  
154 new contacts and contact loss changes the stiffness of the granular system. Particle  
155 sliding also changes the shear stiffness. These phenomena introduce a nonlinear  
156 strain dependent behavior of the granular material [O'Sullivan, 2011]. The contact  
157 nonlinearity is mainly effective at smaller strains, while material nonlinearity arises

158 mainly from a change in the number and the fabric of contacts taking place at larger  
159 strains [O'Sullivan, 2011]. Material nonlinearity is recognized to play an important role  
160 in the weakening process of granular materials [Johnson & Jia, 2005]. There is a variety  
161 of contact laws used in DEM [Di Renzo & Di Maio, 2004], however, in order to  
162 capture particle scale nonlinearity, we use the nonlinear Hertzian contact law. In this  
163 particle-particle contact law, the spring stiffness and the coefficient of damping are  
164 function of particle material properties and the overlap between particles [Hertz,  
165 1882; Di Renzo & Di Maio, 2004]. The normal and tangential contact forces are  
166 calculated as follows:

$$167 \quad F_{pn} = -k_{pn}\delta\varepsilon_{pn} + c_{pn}\delta u_{pn}, \quad (3)$$

$$168 \quad F_{pt} = \min \left\{ \left| k_{pt} \int_{t_{c,0}}^t \delta u_{pt} dt + c_{pt} \delta u_{pt} \right|, \mu_c F_{pn} \right\}, \quad (4)$$

169 where  $k_{pn}$  and  $k_{pt}$  are the normal and tangential spring stiffness,  $c_{pn}$  and  $c_{pt}$  are  
170 the normal and tangential damping coefficient,  $\delta\varepsilon_{pn}$  is the overlap and  $\delta u_{pn}$  and  
171  $\delta u_{pt}$  are the relative normal and tangential velocities of two particles in contact,  
172 respectively. In Eq. (4), the parameter  $\mu_c$  represents the inter-particle friction  
173 coefficient that limits the tangential force. When the tangential contact force between  
174 two particles in contact reaches this limit, they start sliding against each other. The  
175 integral term in Eq. (4) shows an incremental spring, storing energy based on the  
176 relative elastic tangential deformation of the particle surface starting from the  
177 moment particles touch each other at  $t_{c,0}$ . A damping is added to the spring  
178 component of the tangential force if the Coulomb criterion is not met [Di Renzo &



179 *Di Maio, 2004; Goniva et al., 2012*]. The spring and damping coefficients are calculated  
 180 as follows:

$$181 \quad k_{pn} = \frac{4}{3} Y^* \sqrt{R^* \delta \varepsilon_{pn}}, \quad (5)$$

$$182 \quad k_{pt} = 8 G^* \sqrt{R^* \delta \varepsilon_{pn}}, \quad (6)$$

$$183 \quad c_{pn} = -2 \sqrt{\frac{5}{6}} \times \frac{\ln(r)}{\sqrt{\ln^2(r) + \pi^2}} \times \sqrt{2Y^* \sqrt{R^* \delta \varepsilon_{pn}} m^*}, \quad (7)$$

$$184 \quad c_{pt} = -2 \sqrt{\frac{5}{6}} \times \frac{\ln(r)}{\sqrt{\ln^2(r) + \pi^2}} \times \sqrt{8 G^* \sqrt{R^* \delta \varepsilon_{pn}} m^*}, \quad (8)$$

185 where,  $r$  is the restitution coefficient, and  $Y^*$ ,  $R^*$ ,  $G^*$  and  $m^*$  are the equivalent  
 186 Young's modulus, radius, shear modulus and mass, respectively, calculated as  
 187 follows:

$$188 \quad \frac{1}{Y^*} = \frac{(1-\nu_1^2)}{Y_1} + \frac{(1-\nu_2^2)}{Y_2}, \quad (9)$$

$$189 \quad \frac{1}{G^*} = \frac{2(2-\nu_1)(1+\nu_1)}{Y_1} + \frac{2(2-\nu_2)(1+\nu_2)}{Y_2}, \quad (10)$$

$$190 \quad \frac{1}{R^*} = \frac{1}{R_1} + \frac{1}{R_2}, \quad (11)$$

$$191 \quad \frac{1}{m^*} = \frac{1}{m_1} + \frac{1}{m_2}, \quad (12)$$

192 where subscripts 1 and 2 refer to the two particles in contact and  $\nu$  is the Poisson's  
 193 ratio of the particle.

194 Figure 1 illustrates a granular layer representing a part of a granular fault gouge.  
 195 In our model, 8000 spherical particles constitute this layer with particle diameter  
 196 ranging 90-150  $\mu\text{m}$  having a uniform, poly-disperse particle size distribution. The  
 197 sample size in our simulations is  $11 \times 1.5 \times 0.8 \text{ mm}^3$ . This sample size is large enough

198 to show proper effect of 3D particle interaction [Ferdowsi, 2014] as well as  
199 jamming/unjamming transitions in the granular layer [Marone et al., 2008]. On the  
200 sample top and bottom, we employ two corrugated plates with high surface  
201 roughness modeled by a friction coefficient of 0.9 between the plates and particles  
202 to facilitate the transmission of shear stresses to the granular gouge. This kind of  
203 geometry is inspired from the BIAx laboratory earthquake machine using corrugated  
204 driving blocks [Marone, 1998a; Rivière et al., 2018]. In this paper, our study is dedicated  
205 to influence of the particle roughness of fault gouge material represented by inter-  
206 particle friction coefficient on stick-slip behavior. We do not study the effect of  
207 characteristics of the corrugated plates e.g. length, depth, etc. on the dynamics of the  
208 sheared granular layer, since a detailed analysis showed that change of these  
209 characteristics does not alter the dynamic regime. On the front- and back-sides of  
210 the sample, we implement frictionless walls with the same elastic properties of  
211 particles. This type of interaction between particles and walls are designed to avoid  
212 rigid wall boundary conditions. Wall-particle interaction in our DEM model is the  
213 same as particle-particle interaction when one particle has an infinite radius. Periodic  
214 boundary conditions are applied at the left and right sidewalls representing a long  
215 fault gouge in  $x$  direction. The periodic boundary conditions allow for large shear  
216 displacements, and facilitate recording many slip events to be used for statistical  
217 analyses. To prepare the sample, particles are inserted randomly in space descending  
218 with an initial velocity of  $10^{-2}$  cm/s. Next, the upper plate is moved downward to  
219 apply a confining stress to confine the sample. At this stage, the confining stress

220 increases until the desired confining stress is attained (10 MPa). The position of the  
221 upper plate is adapted continuously, as in the lab experiments, in order to maintain  
222 the confining stress constant. At constant confining stress, shearing is initiated by  
223 moving the bottom plate in  $x$  direction with a displacement-controlled mechanism  
224 (constant velocity of 600  $\mu\text{m/s}$ ) until reaching the maximum shear stress, at which  
225 point the stick-slip process commences. The particle density is 2900  $\text{kg/m}^3$  that  
226 results in an applied time step for DEM calculations of  $15 \times 10^{-9}$  seconds, within the  
227 recommended range based on the Rayleigh time. Our DEM calculations remain in  
228 the quasi-static regime by controlling the inertial number to be below  $10^{-3}$  [*MiDi*,  
229 *2004*; *Sheng et al., 2004*; *Agnolin & Roux, 2007*]. Similar to our previous works  
230 [*Dorostkar et al., 2017c*; *Dorostkar et al., 2018*], we use LIGGGHTS [*Goniva et al., 2012*;  
231 *Kloss et al., 2012*] to model the granular fault gouge.

### 232 **3- Results**

233 We show in Fig. 2a and 2b the evolution of macroscopic friction and gouge thickness,  
234 respectively, for three different inter-particle frictions of 0.1, 0.5 and 0.9 during the  
235 stick-slip dynamics. The macroscopic friction is defined as the ratio of shear stress  
236 to normal stress (confining stress) on the driving block and the gouge thickness  
237 represents the measurement in  $z$  direction of Fig. 1. Figure 2c shows with higher  
238 resolution the shaded area in Fig. 2a. The macroscopic friction (or shear stress since  
239 the confining stress is kept constant) increases nonlinearly reaching to a critical state  
240 where micro-slips take place followed by a major slip event. The average macroscopic  
241 friction increases nonlinearly with the inter-particle friction coefficient: at low inter-

242 particle coefficients (0.1 to 0.5), the macroscopic friction coefficient increases much  
243 more compared to high inter-particle coefficients (0.5 and 0.9). A similar behavior  
244 can be observed for the gouge thickness (Fig. 2b).

245 We perform long simulations for inter-particle friction coefficients between 0.1 -  
246 1 and measure the average macroscopic friction over all stick and slip phases (Fig.  
247 3a). We observe that the gouge strength represented by macroscopic friction  
248 increases nonlinearly with inter-particle friction and saturates at around  $\mu_c = 0.9$  to  
249 1. A similar behavior is observed for gouge thickness (Fig. 3b). The standard  
250 deviations in macroscopic friction signal and gouge thickness are found to increase  
251 with increasing inter-particle friction coefficient (Fig. 3a and 3b). We calculate the  
252 slip recurrence time for slip events with a drop in macroscopic friction larger than  
253 0.01, a threshold that avoids capturing micro-slips before a major slip event  
254 [Dorostkar *et al.*, 2017b]. The slip recurrence time and its standard deviation decreases  
255 with increasing inter-particle friction coefficient, meaning that slip events occur more  
256 often and more regularly in a gouge with higher frictional particles (Fig. 3c)

257 To complement the observations in Fig. 3, in Fig. 4a, we observe an almost linear  
258 relation between macroscopic friction and gouge thickness for simulations with  
259 different inter-particle friction coefficient. The more dense population of data points  
260 at higher thickness and macroscopic friction is consistent with the nonlinear behavior  
261 observed in Figs. 2b and 2c. We also observe that the number of slip events increases  
262 with increasing inter-particle coefficient (equivalent to a decreasing recurrence time  
263 as seen in figure 3c) The histogram of slip events' friction drop (Fig. 5) shows that

264 this increase in number of events at higher inter-particle friction coefficient mainly  
265 stems from smaller events: a fault gouge with higher inter-particle friction coefficient  
266 experiences more smaller slip events at a shorter and more regular inter-event time  
267 (less standard deviation in Fig. 3c) with lower friction (stress) drop. Remark however  
268 that, at high inter-particle friction, although the majority of the slip events are smaller,  
269 there are also some extreme slip events, which are larger than the events at lower  
270 inter-particle friction.

271 When studying the time evolution of the average contact force, we interestingly  
272 find a similar stick-slip type of behavior (Fig. 6). This observation shows that the  
273 macroscopic response of the sheared granular gouge is controlled by the contact  
274 forces at grain scale, i.e. by its grain scale behavior. We observe that, by increasing  
275 the inter-particle friction coefficient, the average contact force increases, however  
276 the relative increase of higher inter-particle friction coefficient becomes smaller at  
277 higher values. The decomposition of contact force into normal and tangential  
278 components shows that the contribution of the normal component is dominant  
279 irrespective of inter-particle friction coefficient. The relative contribution of normal  
280 contact force to the total contact force compared to the contribution of the tangential  
281 contact force decreases for higher inter-particle frictions. This means that, despite a  
282 small increase in tangential contact force, the overall enhanced shear strength of fault  
283 gouge (higher macroscopic friction) at higher inter-particle friction coefficient cannot  
284 solely stem from the increase of tangential contact force. In other words, although  
285 the inter-particle friction coefficient is not directly contributing to the normal contact

286 force, its increase results in a higher normal contact force enhancing the shear  
287 capacity of the system. The hypothesis is that, there exists a structural effect where  
288 at higher inter-particle coefficient less slipping contacts occur, providing a better  
289 support for a contact network to build up higher contact forces. Indeed, when  
290 studying the slipping contact ratio (SCR), which is defined as the ratio between  
291 number of contacts at Coulomb frictional limit prone to slip and the total number  
292 of contacts, we observe drops in SCR of almost one order of magnitude when  
293 changing the inter-particle friction from 0.1 to 0.5 or from 0.5 to 0.9 (Fig. 7a). At the  
294 same time, Fig. 7b shows that the average coordination number (coordination  
295 number is the number of contacts per particle) decreases with increasing particle  
296 friction coefficient. The decrease in coordination number for higher inter-particle  
297 frictions can be attributed to the larger dilation or higher gouge thickness (see Fig.  
298 2b). We remark that, since the coordination number decreases with increasing inter-  
299 particle friction coefficient (Fig. 7b), we also check the total number of slipping  
300 contacts that is not normalized by the number of contacts and observe similar  
301 behavior: the total number of slipping contacts decreases with increasing particle  
302 friction coefficient.

303 In Figs. 7c and 7d, we show the instantaneous and cumulative particle  
304 displacement averaged over all particles, respectively. The instantaneous particle  
305 displacement shows larger jumps upon micro- or major slips for higher inter-particle  
306 friction coefficients. Moreover, the cumulative particle displacement clearly shows  
307 that for a given instant in time (or a given shear strain), particles with higher inter-

308 particle friction underwent higher total displacements. The cumulative displacement  
309 is the total displacement of a particle from the start of each simulation. Our analysis  
310 shows that particle displacement is mainly in x-direction (99 %) along the moving  
311 boundary that imposes the shear stress (Fig. 1).

312 We also study the evolution of potential and kinetic energies in sheared granular  
313 fault gouge (Fig. 8). The elastic strain potential energy is stored within the particle-  
314 particle contacts through overlap between particles [*Dorostkar & Carmeliet, 2018*] and  
315 the kinetic energy is due to translation and rotation of particles. The fault gouge with  
316 higher inter-particle friction coefficient shows more potential energy, however, the  
317 increase in average potential energy is not linear with the increase of inter-particle  
318 friction. For instance, an increase of friction coefficient by a factor 9, from 0.1 (Fig.  
319 8c) to 0.9 (Fig. 8a), leads to an increase in potential energy only by a factor 3. We also  
320 observe that the kinetic energy signal for higher inter-particle frictions shows more  
321 fluctuations and bursts i.e. indicating important rearrangements of particles inside  
322 the fault gouge. While the potential energy at low inter-particle friction coefficient  
323 (Fig. 8c) shows a plateau before the major slip events, the energy is always increasing  
324 approaching slip events for higher inter-particle frictions (Fig. 8a).

325 We compare in Fig. 9 the evolution of macroscopic friction and gouge thickness  
326 for inter-particle friction coefficients with 2 orders of magnitude contrast i.e. 10, 1  
327 and 0.1. The nucleation phase of slip events for higher inter-particle friction  
328 coefficient (Fig.9a) shows a complex behavior in both macroscopic friction and  
329 gouge thickness signals, where a considerable amount of small drops in friction

330 coefficient occur during the stick phase before an upcoming extreme slip event that  
331 has a long recurrence time. The statistical analysis of the size of all slip events based  
332 on drop in macroscopic friction (Fig. 10) shows that slip events occur more often  
333 for higher inter-particle friction coefficients i.e. larger number of slip events with  
334 shorter recurrence interval. More importantly, we observe that, although the higher  
335 number of slip events stems mainly from smaller events (see also Fig. 5), there exist  
336 also some very large slip (extreme) events (Fig. 10).

337 A commonly used technique to monitor fault activity in experiments is Acoustic  
338 Emission (AE) [McLasky & Glaser, 2011; Goebel et al., 2012; Johnson et al., 2013; Goebel  
339 et al., 2014b; Goebel et al., 2017; Rivière et al., 2018]. The origin of AE during laboratory  
340 stick-slip experiments is still under debate, where the AE is suggested to originate  
341 from groaning, creaking, and chattering of continuous grain motions and breakage  
342 of force chains within the fault gouge [Rouet-Leduc et al., 2017; Rivière et al., 2018]. Our  
343 analyses in this study show that the velocity (or acceleration) signal of a flagged  
344 particle demonstrates a very similar behavior to acoustic data of the lab (Fig. 11a),  
345 containing similar information for the prediction of macroscopic friction signal (e.g.  
346 [Rouet-Leduc et al., 2017]). However, since the velocity profile of a flagged particle  
347 contains both motion from arriving waves owing to rearrangements of other particles  
348 and the motion of the particle itself, we call it “pseudo AE” signal. In other words,  
349 although our AE signal is derived from the motion of a single, flagged particle  
350 compared to the lab signal where the AE is usually reordered with a device outside  
351 of the fault gouge; we find that the pseudo AE signal contains sufficient information



352 for the purpose of this study, where we compare the velocity signal from the same  
353 particles for simulations with different inter-particle friction coefficient. The velocity  
354 signal shows bursts at slip events, which are larger for higher inter-particle  
355 coefficients (Fig. 11b). The complementary Cumulative Distribution Function  
356 (cCDF) of pseudo acoustic emission bursts (for all emissions without threshold,  
357 during both stick and slip phases) in Fig. 12 clearly shows the increase of AE  
358 amplitude for higher inter-particle friction coefficients. We remark that the  
359 observations in Fig. 11 and 12 are not dependent on the flagged particle and are  
360 consistently recorded for several chosen particles.

361 We further look at the micromechanics of fault by visualizing the particles on a  
362 plane representative of the whole box (here the front plane of the gouge) for inter-  
363 particle friction coefficients of 0.1 and 10, at a point during the stick phase (Fig. 13).  
364 For panels a and b, we set the maximum of color bar equal to the shear driving plate  
365 velocity, 0.06 cm/s. While a gradient is observed for inter-particle friction of 0.1 from  
366 top to bottom with particles close to the bottom plate having velocity close to 0.06  
367 cm/s, the particle velocity field shows a more uniformly distributed profile for inter-  
368 particle friction of 10. We will discuss the implications of this observation in Section  
369 4. A comparison between panels c and d shows a larger cumulative displacement for  
370 particles with inter-particle friction of 10, where the maximum of color bar is set to  
371 the maximum cumulative displacement of particles with inter-particle friction of 0.1,  
372 for a better comparison. It is clear from Fig. 13d that a larger portion of the sample  
373 has experienced a large displacement. We remind that the main displacement for

374 particles is in x direction, along the driving plate motion. The spatial distributions in  
375 panels c and d are consistent with temporal evolution in Fig. 7d. The spatial  
376 distributions of coordination number for both inter-particle friction coefficients are  
377 rather uniform, while the gouge with lower inter-particle friction shows higher  
378 coordination numbers consistent with observations of Fig. 7b. The gouge with  
379 higher inter-particle friction shows on average lower numbers, also showing some  
380 'blue' spots with low coordination number.

## 381 **4- Discussion**

382 We discuss our main findings and present them in the frame of comparison with  
383 other recent studies on the impact of inter-particle friction coefficient as well as fault  
384 surface roughness. The maximum macroscopic friction coefficient (or frictional  
385 strength of the sheared granular fault gouge) during stick-slip dynamics increases  
386 with inter-particle friction coefficient representing the particle roughness, but  
387 saturates at inter-particle friction values around 0.9-1 (Fig. 3). This observation is  
388 similar to previous studies on peak shear strength of granular materials as well as  
389 stable sliding of a sheared granular layer, where the nonlinear macroscopic behavior  
390 and saturation of friction are observed and attributed to rotation of particles [J.  
391 *Latham et al., 2005; Azéma et al., 2012; Shojaee et al., 2012; Göncü & Luding, 2013;*  
392 *Rathbun et al., 2013; Azéma et al., 2017*]. Knuth and Marone [2007] also observed a  
393 systematic relationship between the strength of granular layers and the surface  
394 roughness of particles [Knuth & Marone, 2007]. Rough faults are suggested to  
395 undergo more deformation in the fault zone requiring more overall work to shear

396 the fault [Rathbun *et al.*, 2013], as we also observe in the higher cumulative particle  
397 displacement for fault gouge with higher particle roughness (Fig. 7 and Fig. 13).  
398 Furthermore, we show a linear relation between macroscopic stick-slip friction and  
399 gouge thickness (Fig. 4), as is confirmed by previous research [Mead, 1925; Marone,  
400 1998b; Frye & Marone, 2002; Knuth & Marone, 2007; Makedonska *et al.*, 2011].

401 The macroscopic frictional strength of a granular layer also follows the bulk  
402 properties (e.g. bulk stiffness) [Rabinowicz, 1956; Knuth & Marone, 2007; Leeman *et al.*,  
403 2016]. It is suggested that particle-scale roughness is responsible for enhancing the  
404 elasticity of a granular system at global scale [Wang *et al.*, 2007]. Similarly, we also  
405 observe in Fig. 6 that the higher particle roughness makes a granular structure  
406 stronger, which is capable of building higher contact forces leading to a higher shear  
407 strength.

408 We observe higher standard deviation for both macroscopic stick-slip friction and  
409 gouge thickness with increasing roughness (Fig. 3). A hypothesis here is that the fault  
410 gouge with higher particle roughness has to undergo more complex states during  
411 deformation towards failure, which involves more locked contacts (and less slipping  
412 contacts; see Fig. 7) in spite of a lower total number of contacts, which makes the  
413 system to experience a wider range of macroscopic frictions and dilations manifested  
414 in higher variability of friction coefficient and gouge thickness (Fig. 3). The system  
415 with rougher particle surface going through those complex topographical states has  
416 to expand more to accommodate the continuous externally applied shear, which  
417 leads to a higher dilatation and a lower coordination number. These conditions make

418 a fault gouge with rougher particle surface to fail more frequently leading to more  
419 fluctuations and variability both in macroscopic friction (slip events) and gouge  
420 thickness. Slip events in a sheared granular layer are reported as a collective  
421 phenomenon: slipping contact ratio increases approaching the failure leading to a  
422 major slip event. Therefore, from another point of view, in a system with a higher  
423 particle roughness and lower slipping contact ratio, the slip event is prevented due  
424 to stronger contacts, so that the system fails only partially leading to a large number  
425 of small slips. We also remark that, since the main displacement for particles is along  
426 the direction of the shear driving plate (x direction in Fig. 1), the higher cumulative  
427 particle displacement is consistent with a higher number of slip events, where at each  
428 slip event there is a displacement (rupture) for the center of mass of the granular  
429 fault gouge.

430 The recent numerical model using 2-D plane strain calculations by Tal and Hager  
431 [2018] showed that, as the roughness amplitude increases, the load in the fault is  
432 released by more slip events but with lower average stress drops [Tal & Hager, 2018].  
433 Similarly, we observe in Fig. 4 that increase of particle roughness leads to more slip  
434 events but with lower friction drops. Using numerical simulations, Tal et al. [2018]  
435 observed a more complex behavior for faults with higher roughness, where the  
436 complexities in the nucleation process are reflected as irregular fluctuations in the  
437 moment rate for rougher faults [Tal et al., 2018]. Interestingly, we observe a similar  
438 complexity in 3D DEM, where the nucleation (stick) phase of slip events contains  
439 many fluctuations i.e. smaller slip events. However, using the advantage of DEM and

440 employing the periodic boundary conditions we can, in contrast to experiments,  
441 shear the fault gouge during long time collecting information of hundreds of slip  
442 events. Using a statistical analysis, we then show that a fault gouge with high enough  
443 particle roughness and large enough shear displacement shows, besides the large  
444 amount of small events, some extreme slip events with long recurrence time  
445 compared to faults with lower roughness.

446 Previous research on continuous monitoring of AE during stick-slip experiments  
447 of rocks has deepened our understanding from the micromechanics of slip events  
448 [McLasky & Glaser, 2011; Goebel et al., 2012]. The experimental work by Kwiatek et  
449 al. [2014] showed that the observed changes in AE characteristics are clearly  
450 correlated to the fault topography and roughness [Kwiatek et al., 2014]. Goebel et al.  
451 [2017] showed that faults with rougher surface show a more spatially distributed AE  
452 activity and a higher b-value [Goebel et al., 2017]. Our DEM simulations also show  
453 dependency of AE on fault particle roughness. We observe higher temporal (Fig. 11  
454 and 12) and more spatially distributed AE (Fig. 13) for higher fault particle  
455 roughness. The higher AE in fault gouge with higher particle surface roughness is  
456 consistent with more frequent kinetic energy releases (Fig. 8), since AE is believed  
457 to originate from the rearrangement of particles. We remind that, although our  
458 discussion on AE is based on velocity tracking of single flagged particles (called here  
459 pseudo AE), the temporal evolution and higher moments of those signals show very  
460 similar behavior to laboratory AE (e.g. [Rouet-Leduc et al., 2017; Rivière et al., 2018]), as  
461 we are using them for machine learning analyses in our ongoing research.

462 Overall, our observations on the effect of particle roughness approximated by inter-  
463 particle friction coefficient using 3-D DEM model of a granular fault gouge show  
464 similarities to other numerical and experimental works on the effect of fault surface  
465 roughness that did not consider the granular gouge. The DEM model in this work is  
466 a vast simplification of real faults in nature but yet expands our understanding of  
467 micro-scale fault roughness and provides a mean to study and measure quantities  
468 that are not feasible to measure in the lab and in the field, showing how numerical  
469 models can boost our understanding from physical process that dictate frictional  
470 strength of a fault damage zone.

## 471 **5- Conclusions**

472 We model stick-slip dynamics of a granular fault gouge by 3-D discrete element  
473 simulations for different values of the micro-scale fault roughness (referred to as  
474 particle roughness) approximated by the inter-particle friction coefficient to better  
475 understand its effect on the characteristics of seismic cycles. The major findings of  
476 this study can be summarized as follows:

- 477 - The fault gouge frictional strength, dilation and their standard deviation  
478 nonlinearly increase with the particle roughness, saturating at high inter-  
479 particle friction.
- 480 - The average slip events' recurrence time and its standard deviation decreases  
481 with particle roughness, meaning that, rougher faults fail more frequently. A  
482 rougher fault shows a more complex nucleation (stick) phase, characterized

483 by many small slip events as manifested by the more frequent energy release  
484 in the kinetic energy signal.

485 - Our statistical analyses on a large number of slip events obtained by shearing  
486 the fault gouge to a large shear strain show that rougher faults show a higher  
487 number of slip events mainly consisting in small slip events (small drops in  
488 friction coefficient), however there are also some extreme slip events larger  
489 than the extreme events of smoother faults.

490 - The fault gouge with higher particle roughness shows higher stored potential  
491 energy and stronger particle-particle contacts, a structure that needs more  
492 work to deform it and therefore, for a given shear strain, the particles  
493 experience more deformation.

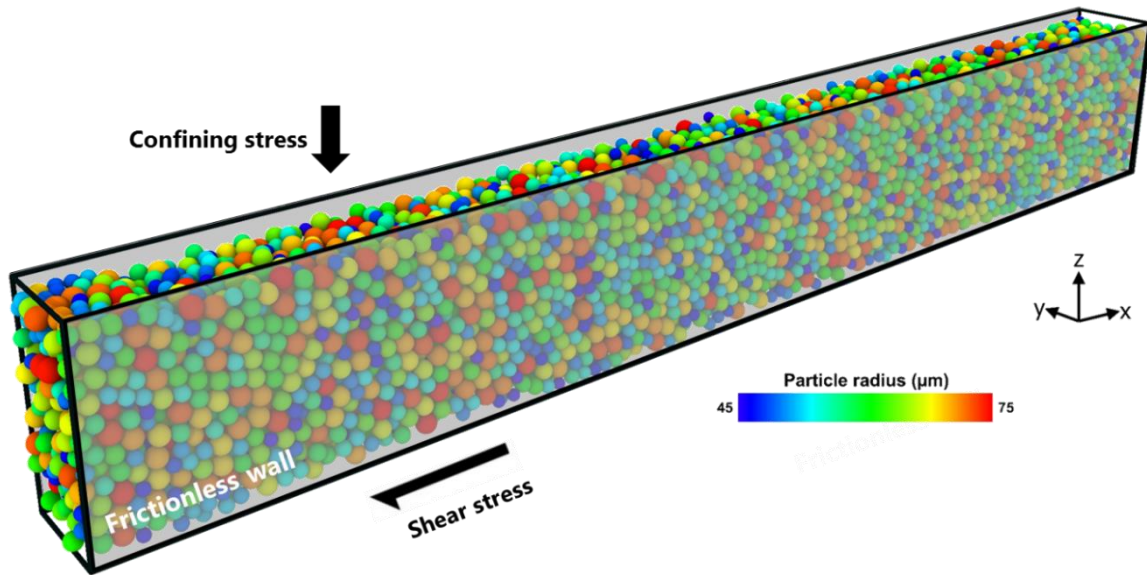
494 - The pseudo acoustic emission analysis, based on the monitoring of the  
495 velocity of particles, shows higher temporal AE for fault gouge with higher  
496 particle surface roughness.

## 497 **6- Acknowledgement**

498 The authors thank ETH Zurich for funding this study. The data related to this paper  
499 can be obtained by contacting the corresponding author at [domid@ethz.ch](mailto:domid@ethz.ch).

500 7- Figures

501



502

503

504

505 **Fig. 1:** Three dimensional granular fault gouge with 8000 particles with poly-disperse  
506 diameter distribution of 90-150 micrometer. The fault gouge is confined in z  
507 direction and sheared in x direction with periodic boundary conditions (Image  
508 produced with the open source visualization tool (OVITO) [Stukowski, 2010]).

509

510

511

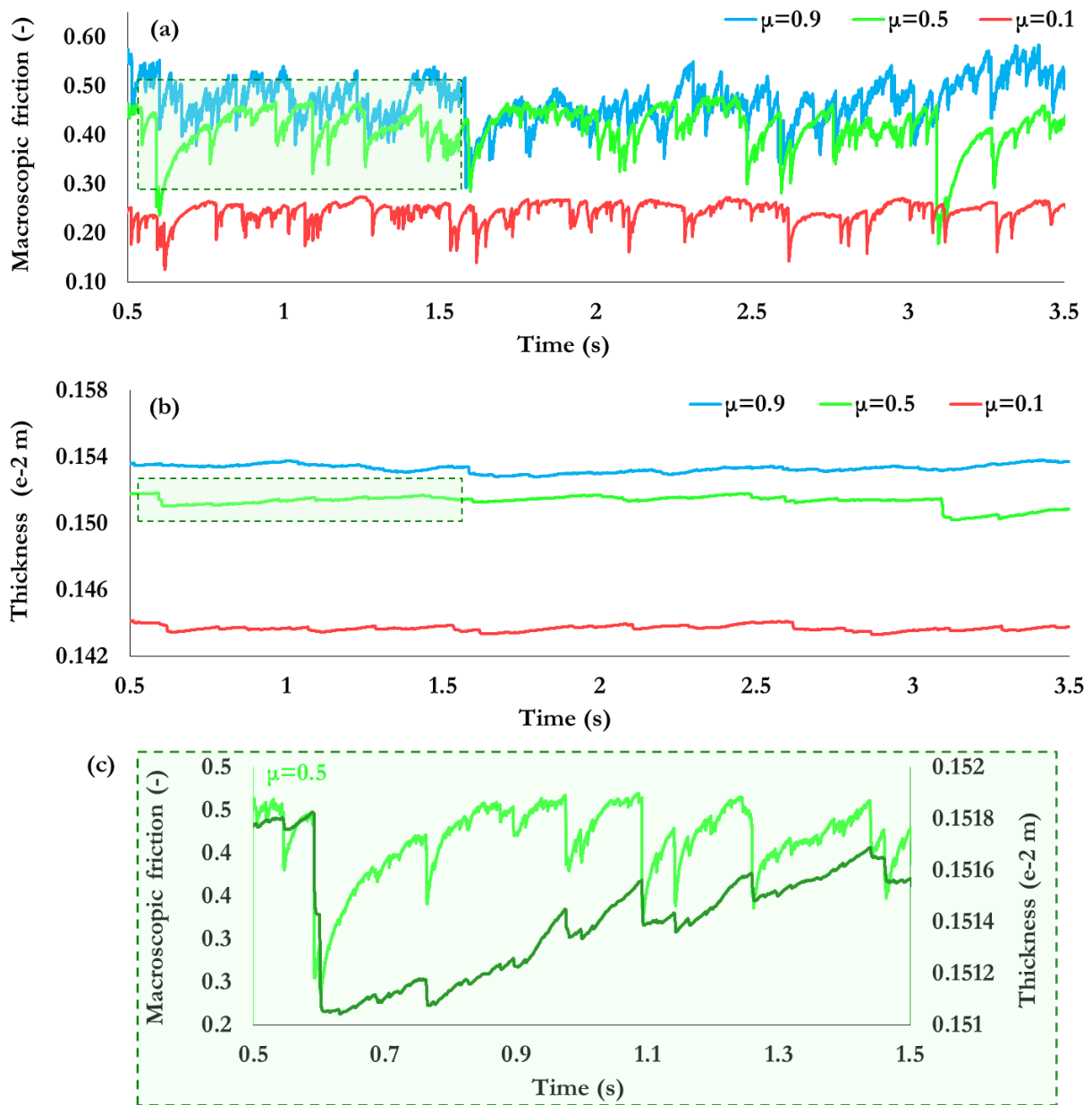
512

513

514

515





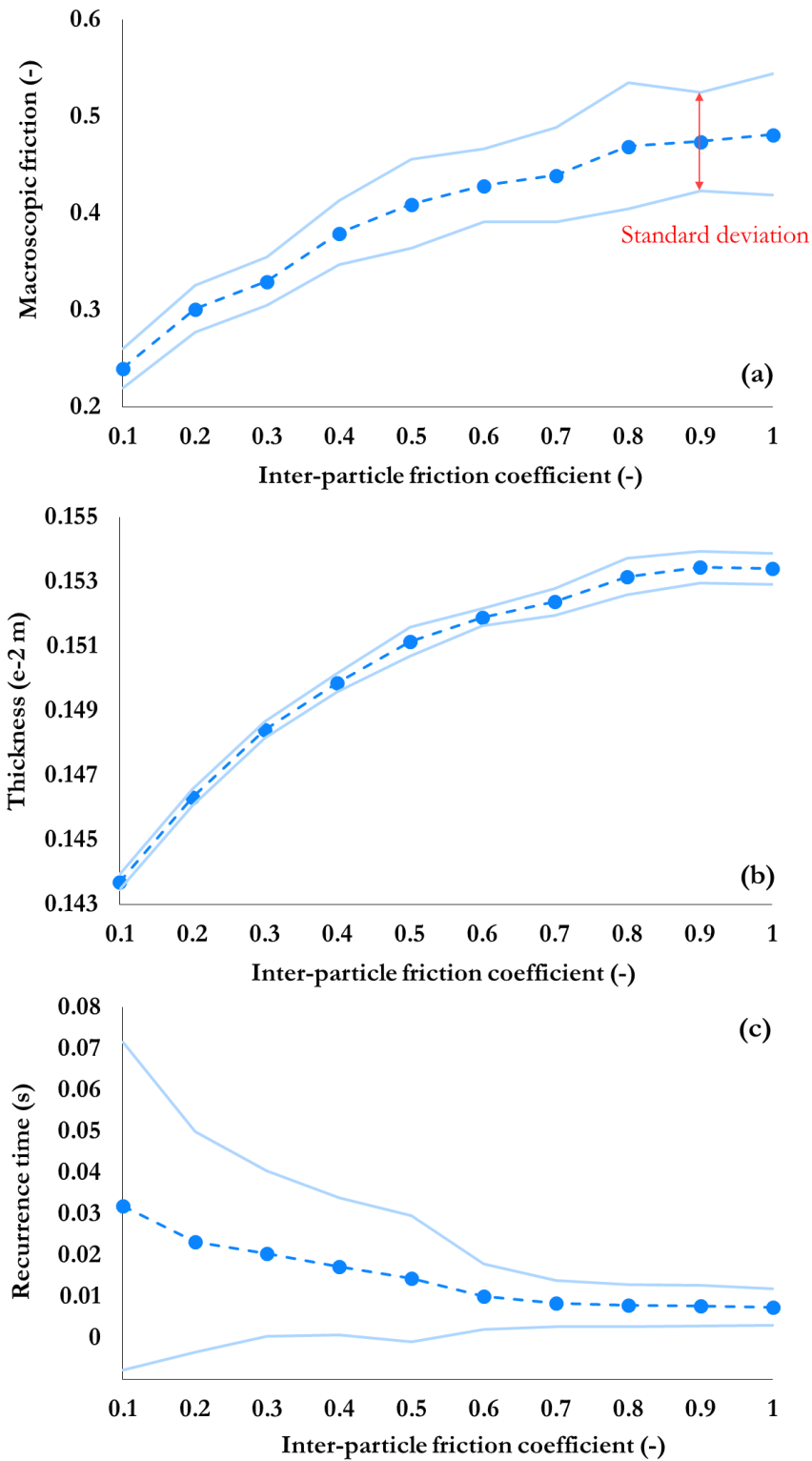
516

517 **Fig. 2:** Time series of (a) macroscopic friction and (b) fault gouge thickness for three  
 518 different inter-particle friction coefficients. The greenish shaded areas in a and b are  
 519 shown in (c) with a higher resolution for  $\mu = 0.5$ . Please note that the thickness is  
 520 shown with a secondary axis in c.

521

522

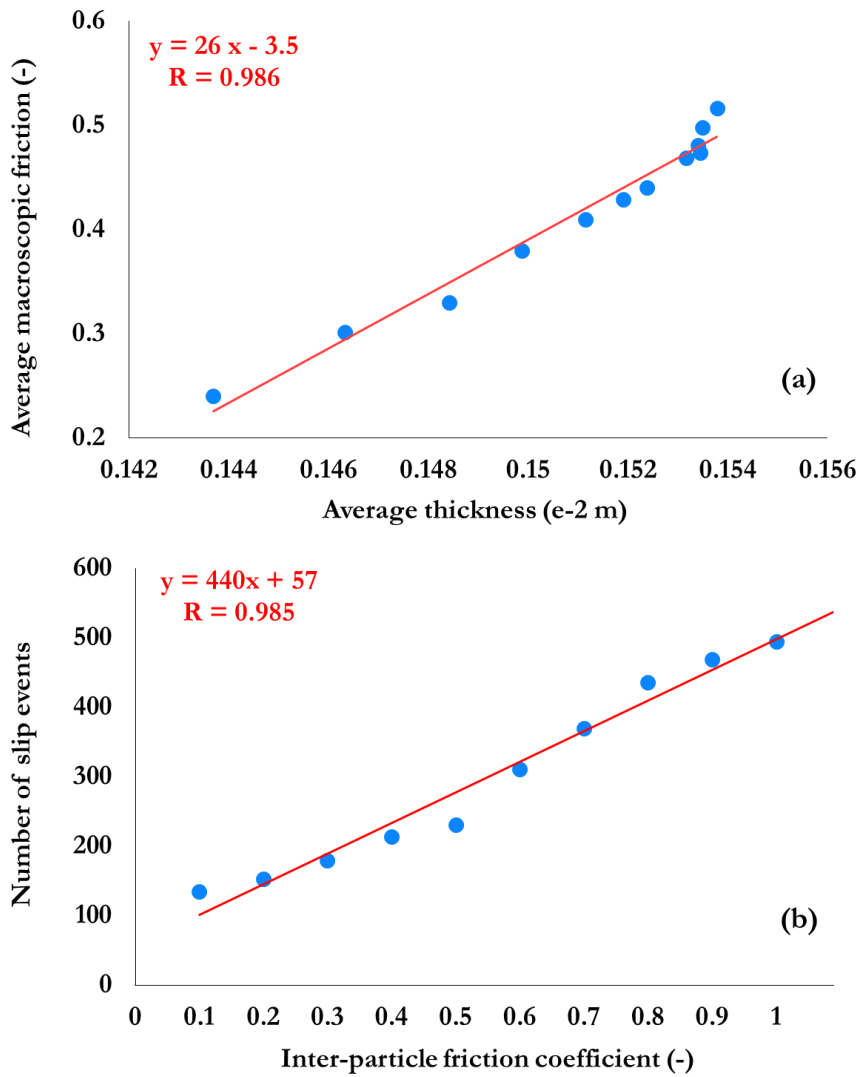
523



524

525

526 **Fig. 3:** Average (a) macroscopic friction, (b) fault gouge thickness and (c) slip  
 527 recurrence time for long time-train stick-slip dynamics as function of the inter-  
 528 particle friction coefficient. The light blue line in each image shows the uncertainty  
 529 limits based on the standard deviation of data.

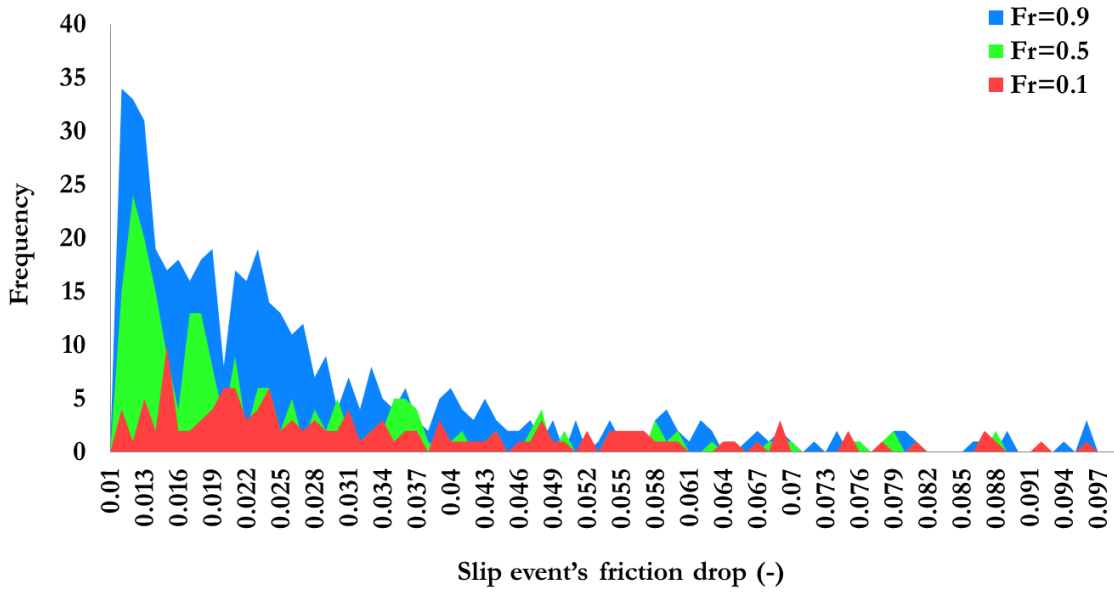


530

531 **Fig. 4:** (a) Average macroscopic friction versus average thickness for stick-slip  
 532 dynamics with different inter-particle friction coefficients. (b) Number of slip events  
 533 versus inter-particle friction coefficient.

534

535

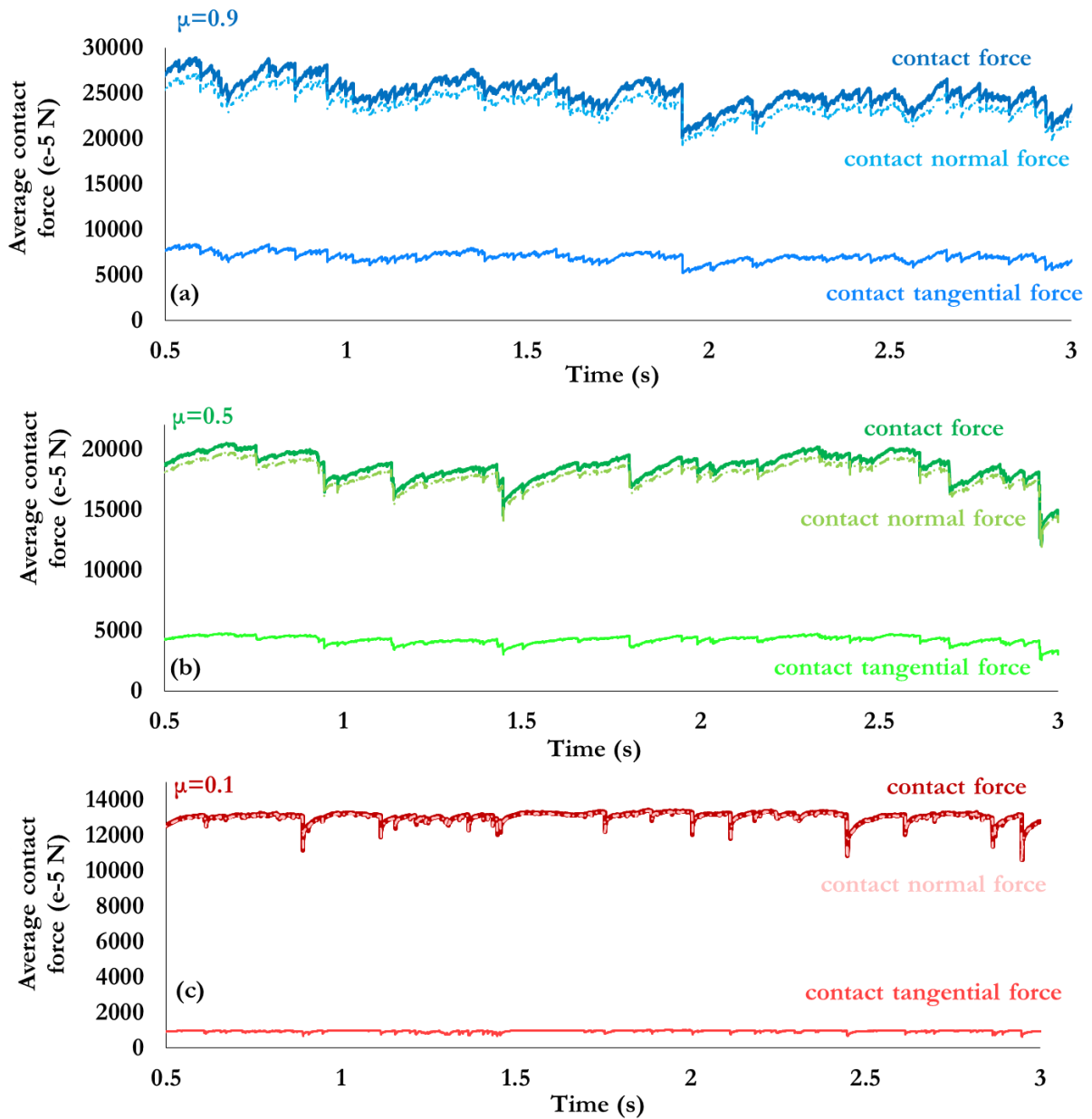


536

537

538 **Fig. 5:** Histogram of slip event's friction drop for three different inter-particle  
 539 friction coefficients. The maximum slip event's friction drop in this histogram is  
 540 limited to 0.1 to highlight and better show the frequency of slip events with small  
 541 friction drop.

542



543

544 **Fig. 6:** (a-c) Average contact force for inter-particle friction coefficients of 0.1, 0.5  
 545 and 0.9, respectively. In each panel, the components of contact force (normal contact  
 546 force and tangential contact force) are separately shown.

547

548

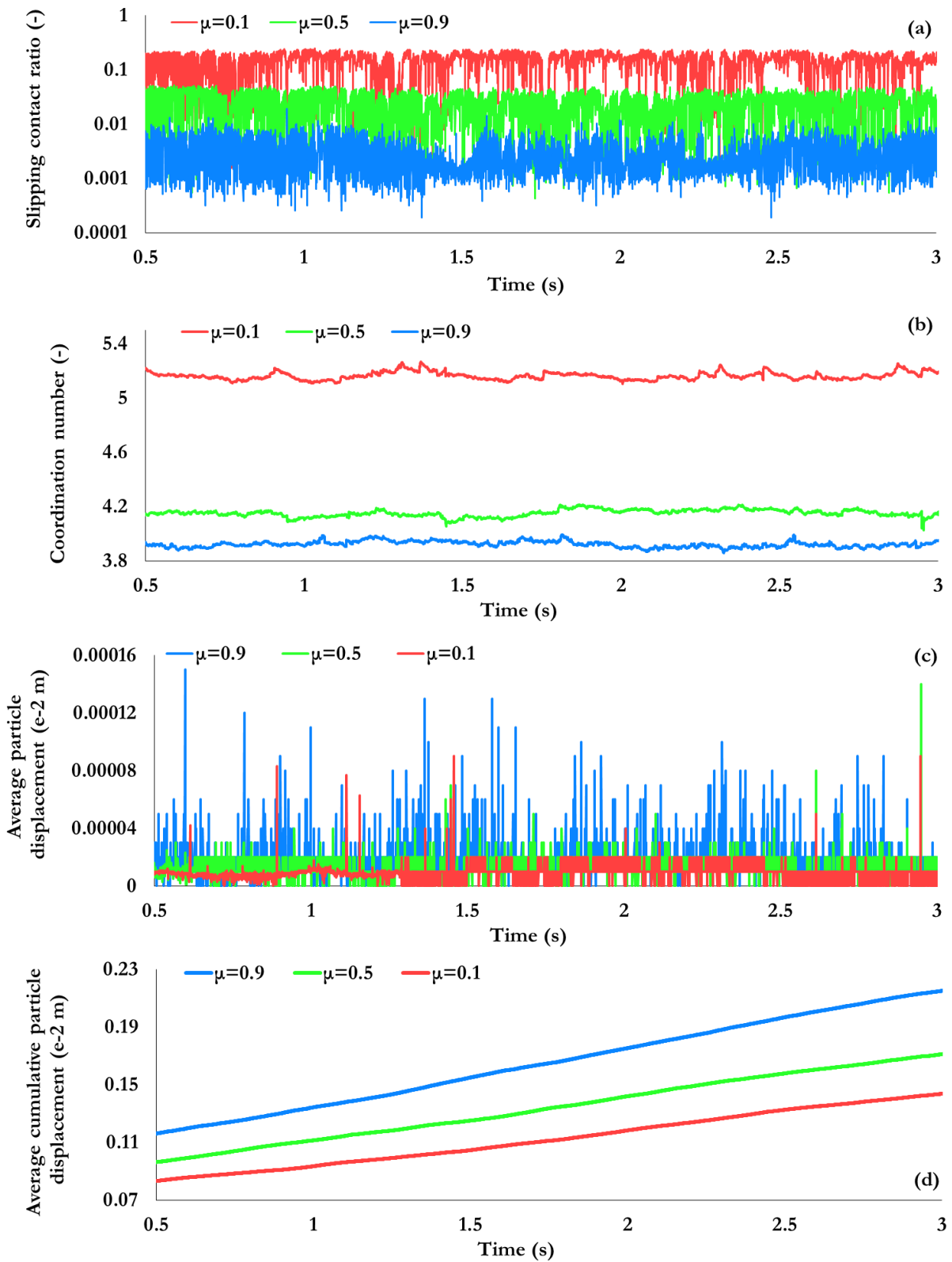
549

550

551

552

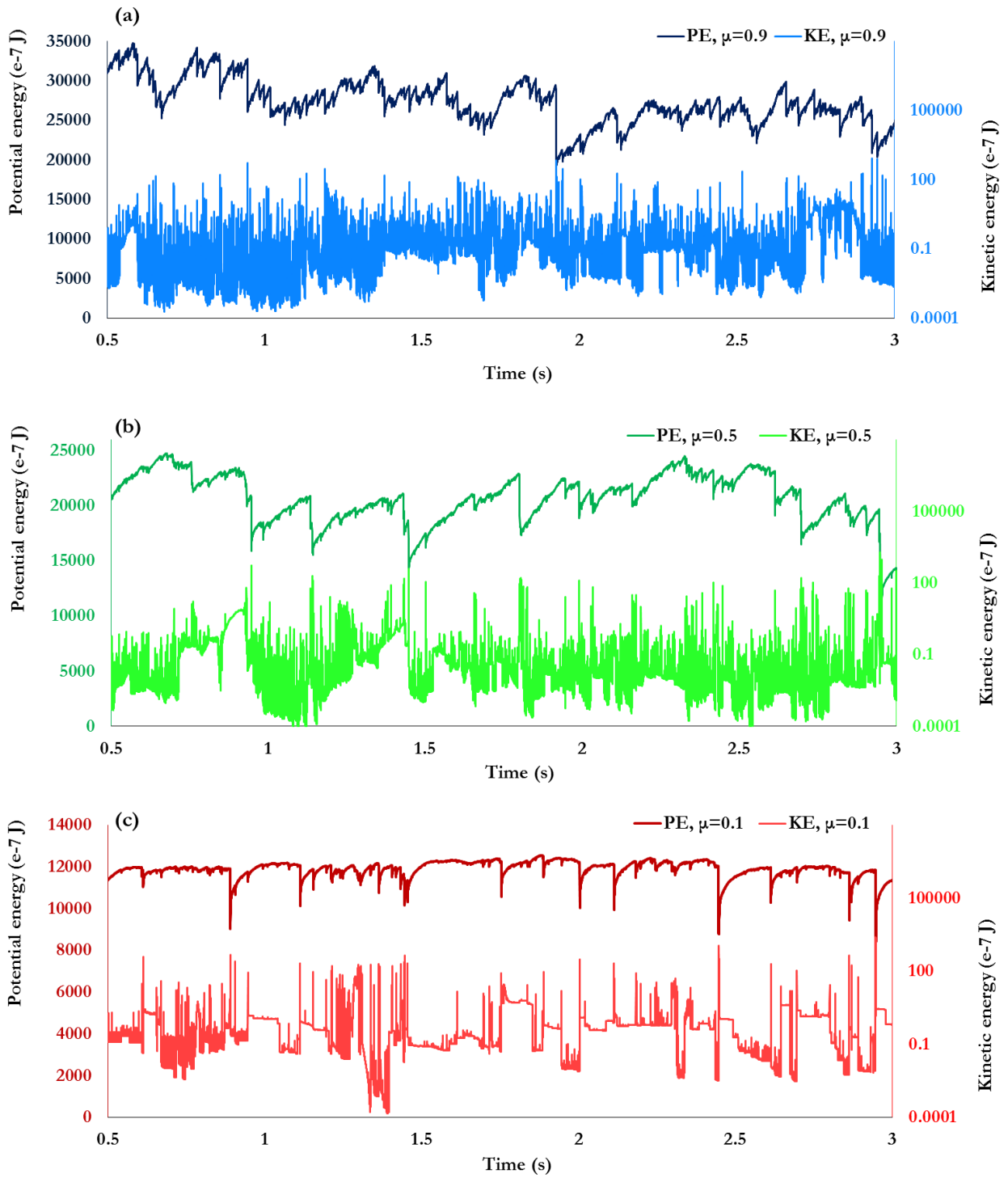
553



554

555 **Fig. 7:** Time series of (a) slipping contact ratio, (b) coordination number, (c) average  
 556 particle displacement and (c) average cumulative particle displacement for inter-  
 557 particle friction coefficients of 0.1, 0.5 and 0.9.

558

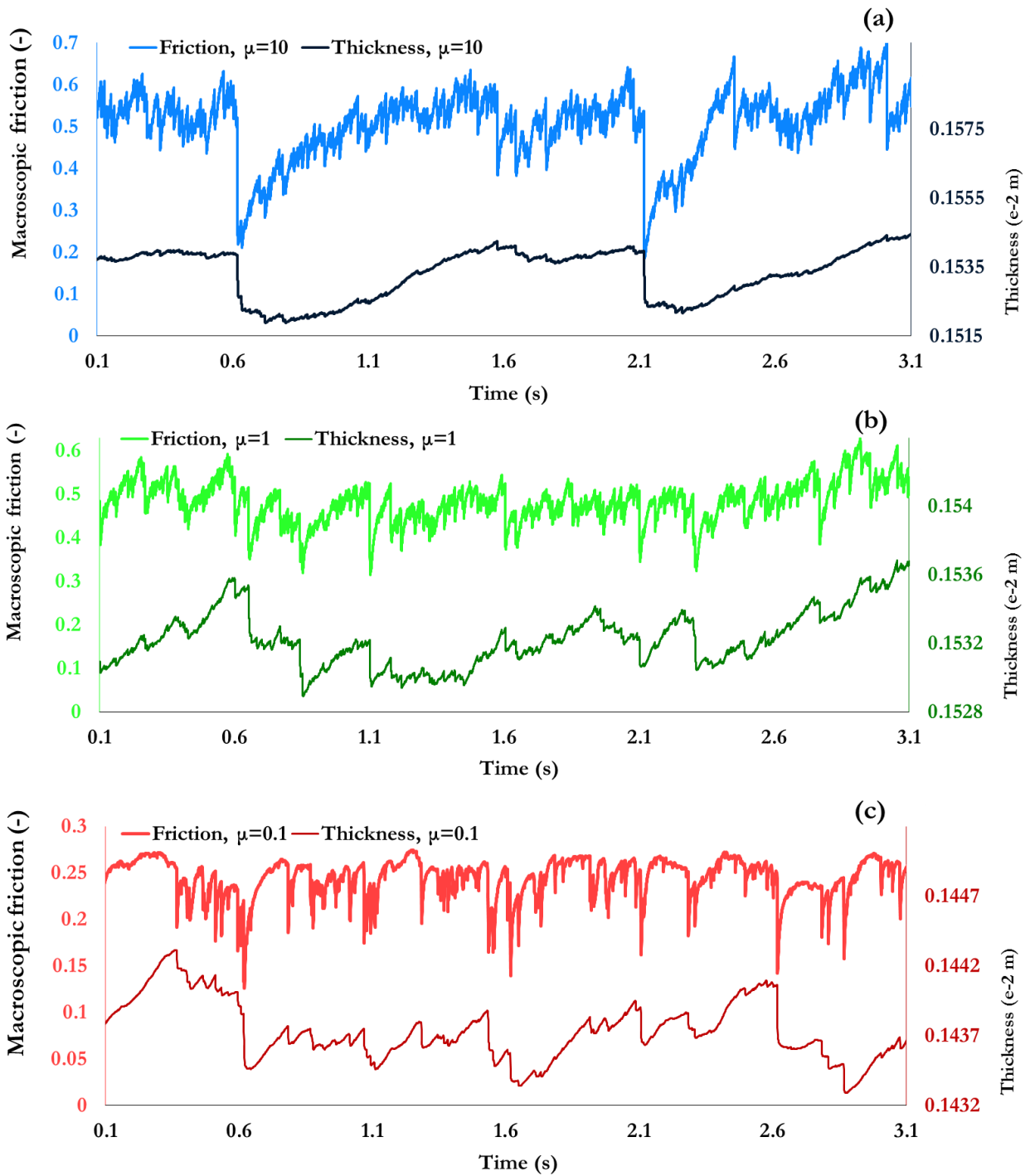


559

560 **Fig. 8:** (a-c) Evolution of potential energy (primary axis, left) and kinetic energy  
 561 (secondary logarithmic axis, right) for inter-particle friction coefficients of 0.9, 0.5  
 562 and 0.1, respectively.

563

564



565

566

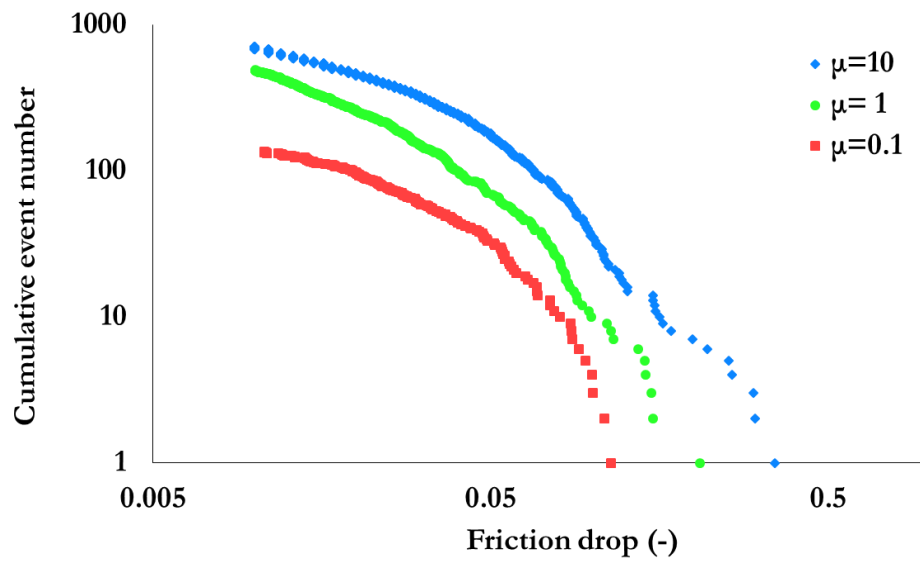
567 **Fig. 9:** (a-c) Time series of macroscopic friction (primary axis, left) and gouge  
 568 thickness (secondary axis, right) for inter-particle friction coefficients of 10, 1, 0.1  
 569 respectively.

570

571

572





573

574

575 **Fig. 10:** Cumulative number of slip events versus friction coefficient drop for  
 576 different inter-particle friction coefficients of 10, 1 and 0.1.

577

578

579

580

581

582

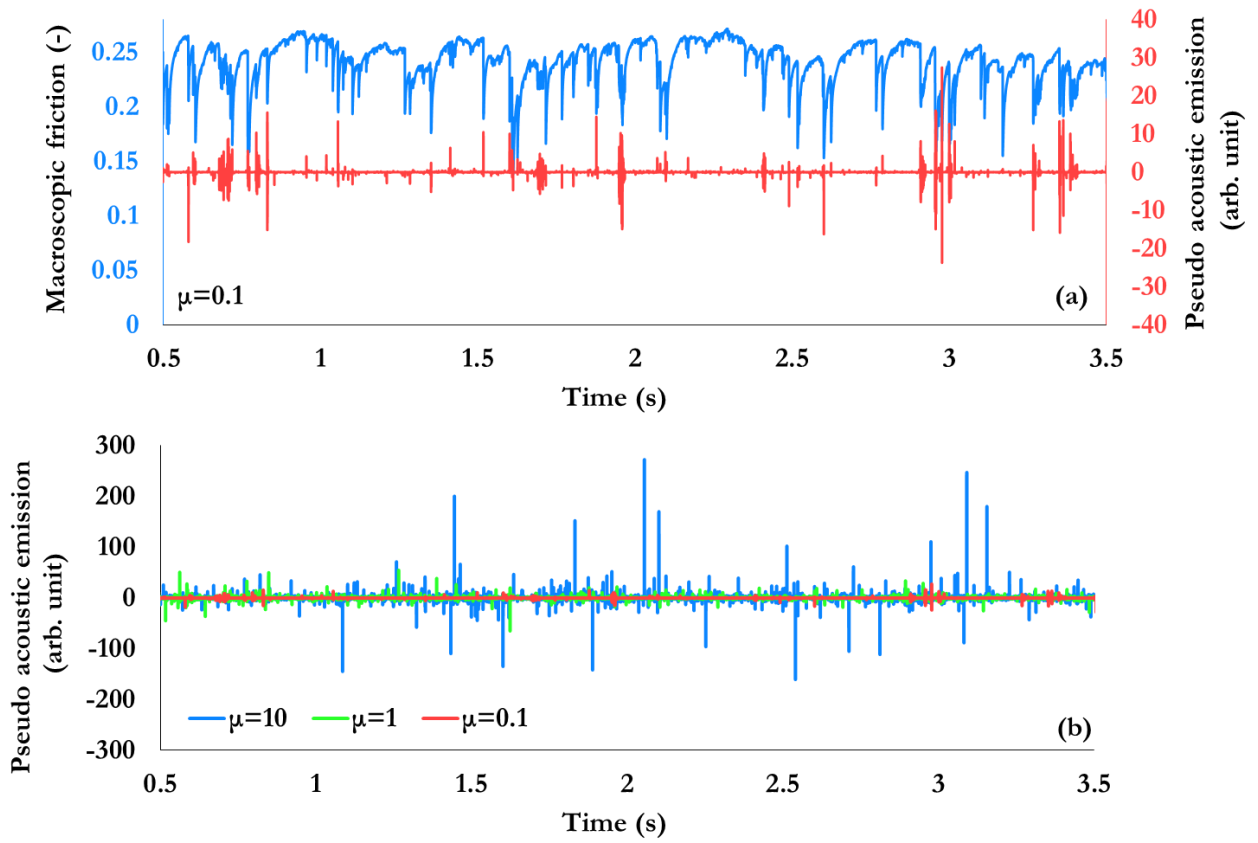
583

584

585

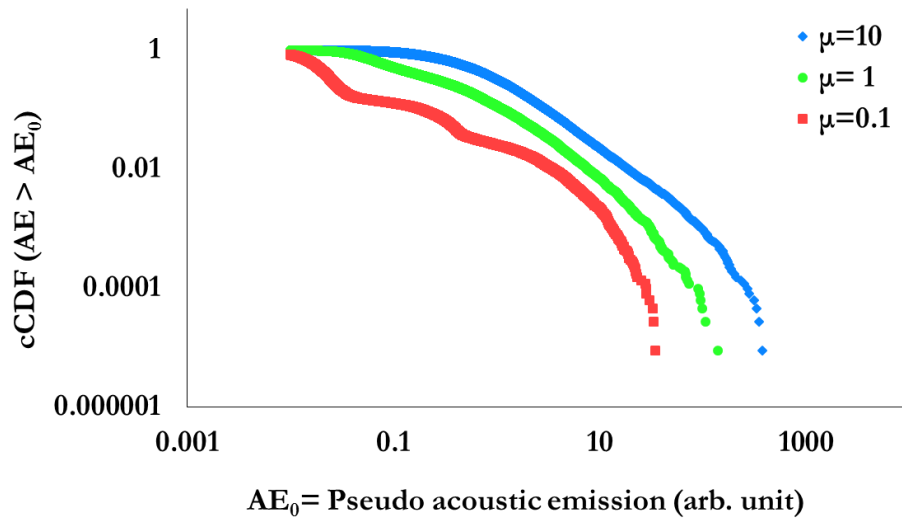
586

587



588  
 589  
 590  
 591  
 592  
 593  
 594  
 595  
 596  
 597  
 598  
 599  
 600  
 601  
 602

**Fig. 11:** (a) Friction coefficient (primary axis, left) and pseudo acoustic emission (secondary axis, right) for inter-particle friction coefficients of 10, 1 and 0.1. (b) Time series of pseudo acoustic emission for inter-particle friction coefficients of 10, 1 and 0.1.



603

604 **Fig. 12:** Complementary Cumulative Distribution Function of pseudo acoustic  
 605 emission for different inter-particle friction coefficients of 10, 1 and 0.1.

606

607

608

609

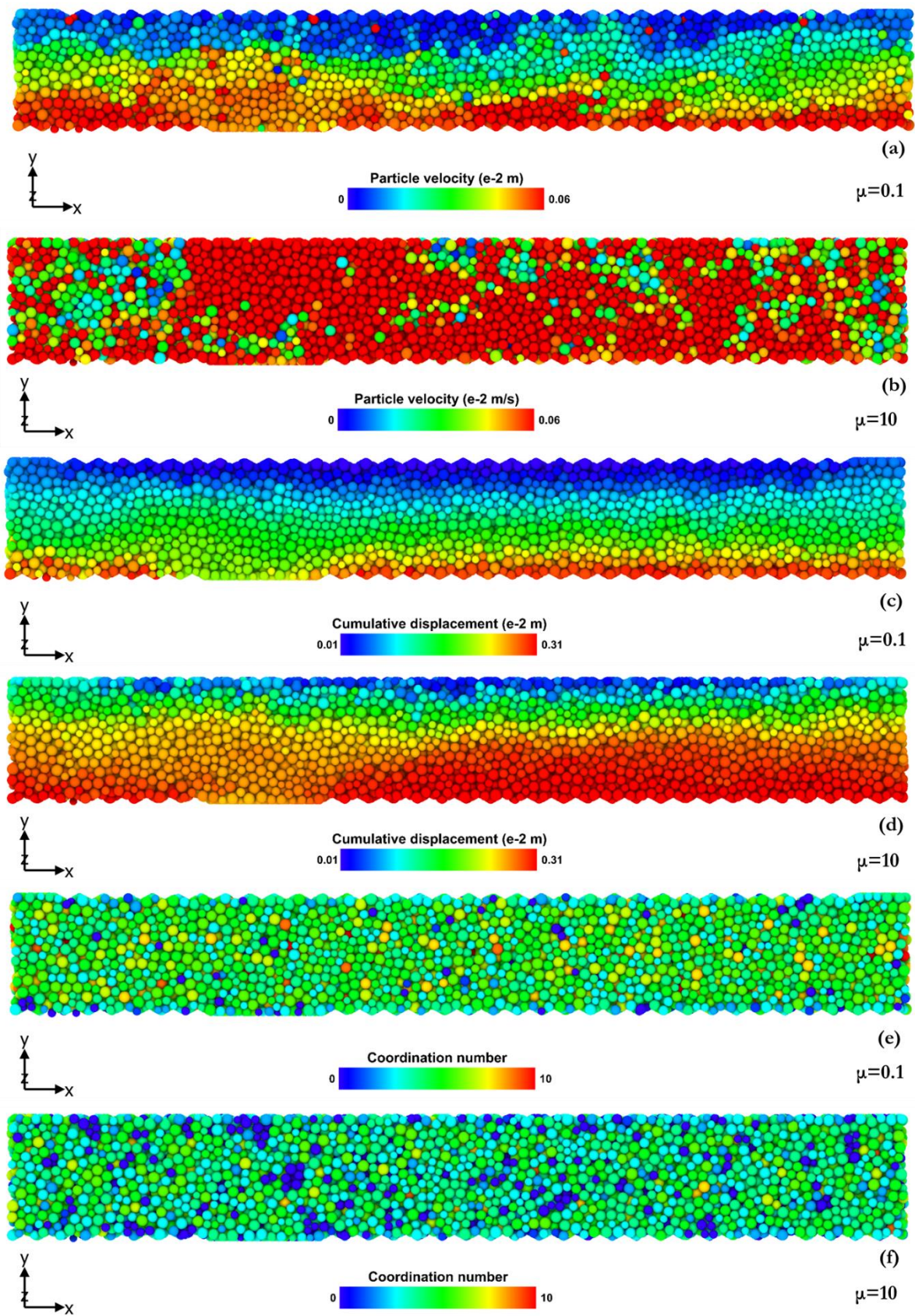
610

611

612

613

614



615

616 **Fig. 13:** (a, c, e) Spatial distribution of particle velocity, cumulative displacement and  
 617 coordination number for inter-particle friction coefficients of 0.1 and (b, d, f) for  
 618 inter-particle friction coefficient of 10, respectively.

## 8- References

- 619  
620  
621 Agnolin, I., & Roux, J. N. (2007). Internal states of model isotropic granular packings.  
622 I. Assembling process, geometry, and contact networks. *Phys. Rev. E*, 76(6).  
623 doi:10.1103/PhysRevE.76.061302
- 624 Angheluta, L., Candela, T., Mathiesen, J., & Renard, F. (2011). Effect of surface  
625 morphology on the dissipation during shear and slip along a rock–rock  
626 interface that contains a visco-elastic core. *Pure Appl. Geophys.*, 168(12), 2335-  
627 2344. doi:10.1007/s00024-011-0272-8
- 628 Anthony, J. L., & Marone, C. (2005). Influence of particle characteristics on granular  
629 friction. *J. Geophys. Res.*, 110(B8). doi:10.1029/2004jb003399
- 630 Azéma, E., Estrada, N., Preechawuttipong, I., Delenne, J.-Y., & Radjai, F. (2017).  
631 Systematic description of the effect of particle shape on the strength  
632 properties of granular media. *EPJ Web Conf.*, 140, 06026.
- 633 Azéma, E., Estrada, N., & Radjaï, F. (2012). Nonlinear effects of particle shape  
634 angularity in sheared granular media. *Phys. Rev. E*, 86(4), 041301.  
635 doi:10.1103/PhysRevE.86.041301
- 636 Biegel, R. L., Sammis, C. G., & Dieterich, J. H. (1989). The frictional properties of a  
637 simulated gouge having a fractal particle distribution. *J. Struct. Geol.*, 11(7), 827-  
638 846. doi:https://doi.org/10.1016/0191-8141(89)90101-6
- 639 Bistacchi, A., Griffith, W. A., Smith, S. A. F., Di Toro, G., Jones, R., & Nielsen, S.  
640 (2011). Fault roughness at seismogenic depths from lidar and  
641 photogrammetric analysis. *Pure Appl. Geophys.*, 168(12), 2345-2363.  
642 doi:10.1007/s00024-011-0301-7
- 643 Bouchon, M., Karabulut, H., Bouin, M.-P., Schmittbuhl, J., Vallée, M., Archuleta, R.,  
644 . . . Marsan, D. (2010). Faulting characteristics of supershear earthquakes.  
645 *Tectonophysics*, 493(3), 244-253.  
646 doi:https://doi.org/10.1016/j.tecto.2010.06.011
- 647 Brantut, N., Schubnel, A., Rouzaud, J. N., Brunet, F., & Shimamoto, T. (2008). High-  
648 velocity frictional properties of a clay-bearing fault gouge and implications for  
649 earthquake mechanics. *J. Geophys. Res.*, 113(B10). doi:10.1029/2007JB005551

- 650 Brodsky, E. E., Gilchrist, J. J., Sagy, A., & Collettini, C. (2011). Faults smooth  
651 gradually as a function of slip. *Earth Planet. Sci. Lett.*, *302*(1), 185-193.  
652 doi:<https://doi.org/10.1016/j.epsl.2010.12.010>
- 653 Brodsky, E. E., Kirkpatrick, J. D., & Candela, T. (2016). Constraints from fault  
654 roughness on the scale-dependent strength of rocks. *Geology*, *44*(1), 19-22.
- 655 Bruhat, L., Fang, Z., & Dunham, E. M. (2016). Rupture complexity and the  
656 supershear transition on rough faults. *J. Geophys. Res.*, *121*(1), 210-224.  
657 doi:[doi:10.1002/2015JB012512](https://doi.org/10.1002/2015JB012512)
- 658 Candela, T., Renard, F., Bouchon, M., Brouste, A., Marsan, D., Schmittbuhl, J., &  
659 Voisin, C. (2009). Characterization of fault roughness at various scales:  
660 Implications of three-dimensional high resolution topography measurements.  
661 *Pure Appl. Geophys.*, *166*(10), 1817-1851. doi:[10.1007/s00024-009-0521-2](https://doi.org/10.1007/s00024-009-0521-2)
- 662 Candela, T., Renard, F., Bouchon, M., Schmittbuhl, J., & Brodsky, E. E. (2011a).  
663 Stress drop during earthquakes: Effect of fault roughness scaling. *Bull. Seismol.*  
664 *Soc. Am.*, *101*(5), 2369-2387.
- 665 Candela, T., Renard, F., Klinger, Y., Mair, K., Schmittbuhl, J., & Brodsky, E. E.  
666 (2012). Roughness of fault surfaces over nine decades of length scales. *J.*  
667 *Geophys. Res.*, *117*(B8). doi:[doi:10.1029/2011JB009041](https://doi.org/10.1029/2011JB009041)
- 668 Candela, T., Renard, F., Schmittbuhl, J., Bouchon, M., & Brodsky, E. E. (2011b).  
669 Fault slip distribution and fault roughness. *Geophys. J. Int.*, *187*(2), 959-968.  
670 doi:[10.1111/j.1365-246X.2011.05189.x](https://doi.org/10.1111/j.1365-246X.2011.05189.x)
- 671 Cashman, S., & Cashman, K. (2000). Cataclasis and deformation-band formation in  
672 unconsolidated marine terrace sand, humboldt county, california. *Geology*,  
673 *28*(2), 111-114. doi:[10.1130/0091-7613\(2000\)28<111:CADFIU>2.0.CO;2](https://doi.org/10.1130/0091-7613(2000)28<111:CADFIU>2.0.CO;2)
- 674 Cashman, S. M., Baldwin, J. N., Cashman, K. V., Swanson, K., & Crawford, R.  
675 (2007). Microstructures developed by coseismic and aseismic faulting in near-  
676 surface sediments, san andreas fault, california. *Geology*, *35*(7), 611-614.
- 677 Chester, F. M., & Chester, J. S. (1998). Ultracataclasite structure and friction  
678 processes of the punchbowl fault, san andreas system, california. *Tectonophysics*,  
679 *295*(1), 199-221. doi:[https://doi.org/10.1016/S0040-1951\(98\)00121-8](https://doi.org/10.1016/S0040-1951(98)00121-8)
- 680 Chester, F. M., & Chester, J. S. (2000). Stress and deformation along wavy frictional  
681 faults. *J. Geophys. Res.*, *105*(B10), 23421-23430. doi:[doi:10.1029/2000JB900241](https://doi.org/10.1029/2000JB900241)

- 682 Chester, F. M., Friedman, M., & Logan, J. M. (1985). Foliated cataclasites.  
683 *Tectonophysics*, 111(1), 139-146. doi:https://doi.org/10.1016/0040-  
684 1951(85)90071-X
- 685 Chester, F. M., & Logan, J. M. (1986). Implications for mechanical properties of  
686 brittle faults from observations of the punchbowl fault zone, California. *Pure*  
687 *Appl. Geophys.*, 124(1), 79-106. doi:10.1007/bf00875720
- 688 Di Renzo, A., & Di Maio, F. P. (2004). Comparison of contact-force models for the  
689 simulation of collisions in dem-based granular flow codes. *Chem. Eng. Sci.*,  
690 59(3), 525-541. doi:10.1016/j.ces.2003.09.037
- 691 Dieterich, J. H., & Kilgore, B. D. (1994). Direct observation of frictional contacts:  
692 New insights for state-dependent properties. *Pure Appl. Geophys.*, 143(1), 283-  
693 302. doi:10.1007/bf00874332
- 694 Dieterich, J. H., & Smith, D. E. (2009). Nonplanar faults: Mechanics of slip and off-  
695 fault damage. *Pure Appl. Geophys.*, 166(10), 1799-1815. doi:10.1007/s00024-  
696 009-0517-y
- 697 Dorostkar, O. (2018). *Stick-slip dynamics in dry and fluid saturated granular fault gouge*  
698 *investigated by numerical simulations*. (PhD dissertation), PhD dissertation, ETH  
699 Zurich.
- 700 Dorostkar, O., & Carmeliet, J. (2018). Potential energy as metric for understanding  
701 stick-slip dynamics in sheared granular fault gouge: A coupled CFD–DEM study.  
702 *Rock Mech. Rock Eng.* doi:10.1007/s00603-018-1457-6
- 703 Dorostkar, O., Guyer, R. A., Johnson, P. A., Marone, C., & Carmeliet, J. (2017a). On  
704 the micromechanics of slip events in sheared, fluid saturated fault gouge.  
705 *Geophys. Res. Lett.*, 6101–6108. doi:10.1002/2017GL073768
- 706 Dorostkar, O., Guyer, R. A., Johnson, P. A., Marone, C., & Carmeliet, J. (2017b). On  
707 the role of fluids in stick-slip dynamics of saturated granular fault gouge using  
708 a coupled computational fluid dynamics-discrete element approach. *J. Geophys.*  
709 *Res.* doi:10.1002/2017JB014099
- 710 Dorostkar, O., Guyer, R. A., Johnson, P. A., Marone, C., & Carmeliet, J. (2018).  
711 Cohesion-induced stabilization in stick-slip dynamics of weakly wet, sheared  
712 granular fault gouge. *J. Geophys. Res.* doi:10.1002/2017JB015171



- 713 Dorostkar, O., Johnson, P., Guyer, R., Marone, C., & Carmeliet, J. (2017c). Do fluids  
714 modify the stick-slip behavior of sheared granular media? *Poromechanics VI :  
715 Proceedings of the Sixth Biot Conference on Poromechanics, 2017.* , 158-163.  
716 doi:10.1061/9780784480779.019
- 717 Dunham, E. M., Belanger, D., Cong, L., & Kozdon, J. E. (2011). Earthquake ruptures  
718 with strongly rate-weakening friction and off-fault plasticity, part 2: Nonplanar  
719 faults. *Bull. Seismol. Soc. Am.*, *101*(5), 2308-2322.
- 720 Engelder, J. T. (1974). Cataclasis and the generation of fault gouge. *GSA Bulletin*,  
721 *85*(10), 1515-1522. doi:10.1130/0016-  
722 7606(1974)85<1515:CATGOF>2.0.CO;2
- 723 Faulkner, D. R., Lewis, A. C., & Rutter, E. H. (2003). On the internal structure and  
724 mechanics of large strike-slip fault zones: Field observations of the carboneras  
725 fault in southeastern Spain. *Tectonophysics*, *367*(3), 235-251.  
726 doi:https://doi.org/10.1016/S0040-1951(03)00134-3
- 727 Ferdowsi, B. (2014). *Discrete element modeling of triggered slip in faults with granular gouge:  
728 Application to dynamic earthquake triggering.* (PhD dissertation ), ETH Zurich.
- 729 Fournier, T., & Morgan, J. (2012). Insights to slip behavior on rough faults using  
730 discrete element modeling. *Geophys. Res. Lett.*, *39*(12).  
731 doi:doi:10.1029/2012GL051899
- 732 Frye, K. M., & Marone, C. (2002). Effect of humidity on granular friction at room  
733 temperature. *J. Geophys. Res.*, *107*(B11). doi:10.1029/2001jb000654
- 734 Goebel, T. H. W., Becker, T. W., Sammis, C. G., Dresen, G., & Schorlemmer, D.  
735 (2014a). Off-fault damage and acoustic emission distributions during the  
736 evolution of structurally complex faults over series of stick-slip events.  
737 *Geophys. J. Int.*, *197*(3), 1705-1718. doi:10.1093/gji/ggu074
- 738 Goebel, T. H. W., Becker, T. W., Schorlemmer, D., Stanchits, S., Sammis, C.,  
739 Rybacki, E., & Dresen, G. (2012). Identifying fault heterogeneity through  
740 mapping spatial anomalies in acoustic emission statistics. *J. Geophys. Res.*,  
741 *117*(B3). doi:doi:10.1029/2011JB008763
- 742 Goebel, T. H. W., Candela, T., Sammis, C. G., Becker, T. W., Dresen, G., &  
743 Schorlemmer, D. (2014b). Seismic event distributions and off-fault damage  
744 during frictional sliding of saw-cut surfaces with pre-defined roughness.  
745 *Geophys. J. Int.*, *196*(1), 612-625. doi:10.1093/gji/ggt401



- 746 Goebel, T. H. W., Kwiatek, G., Becker, T. W., Brodsky, E. E., & Dresen, G. (2017).  
747 What allows seismic events to grow big?: Insights from b-value and fault  
748 roughness analysis in laboratory stick-slip experiments. *Geology*, *45*(9), 815-818.  
749 doi:10.1130/G39147.1
- 750 Göncü, F., & Luding, S. (2013). Effect of particle friction and polydispersity on the  
751 macroscopic stress–strain relations of granular materials. *Acta Geotech.*, *8*(6),  
752 629-643. doi:10.1007/s11440-013-0258-z
- 753 Goniva, C., Kloss, C., Deen, N. G., Kuipers, J. A. M., & Pirker, S. (2012). Influence  
754 of rolling friction on single spout fluidized bed simulation.  
755 *PARTICUOLOGY*, *10*(5), 582-591. doi:10.1016/j.partic.2012.05.002
- 756 Haines, S., Marone, C., & Saffer, D. (2014). Frictional properties of low-angle normal  
757 fault gouges and implications for low-angle normal fault slip. *Earth Planet. Sci.*  
758 *Lett.*, *408*, 57-65. doi:10.1016/j.epsl.2014.09.034
- 759 Heermance, R., Shipton, Z. K., & Evans, J. P. (2003). Fault structure control on fault  
760 slip and ground motion during the 1999 rupture of the chelungpu fault,  
761 taiwan. *Bull. Seismol. Soc. Am.*, *93*(3), 1034-1050. doi:10.1785/0120010230
- 762 Hertz, H. (1882). Ueber die berührung fester elastischer körper. *Journal für die reine*  
763 *und angewandte Mathematik*, *92*, 156-171.
- 764 Ivković, B., Djurdjanović, M., & Stamenković, D. (2000). *The influence of the contact*  
765 *surface roughness on the static friction coefficient* (Vol. 22).
- 766 J. Latham, S., Abe, S., & R. Mora, P. (2005). *Macroscopic friction response of rotational and*  
767 *non-rotational lattice solid gouge models in 2d and 3d*. Paper presented at the Powders  
768 and Grains 2005, Stuttgart, Germany.
- 769 Johnson, P. A., Ferdowsi, B., Kaproth, B. M., Scuderi, M., Griffa, M., Carmeliet, J., .  
770 . . Marone, C. (2013). Acoustic emission and microslip precursors to stick-slip  
771 failure in sheared granular material. *Geophys. Res. Lett.*, *40*(21), 5627-5631.  
772 doi:10.1002/2013gl057848
- 773 Johnson, P. A., & Jia, X. (2005). Nonlinear dynamics, granular media and dynamic  
774 earthquake triggering. *Nature*, *437*(7060), 871-874. doi:10.1038/nature04015
- 775 Karner, S. L., & Marone, C. (2000). Effects of loading rate and normal stress on  
776 stress drop and stick-slip recurrence interval. *GeoComplexity and the Physics of*  
777 *Earthquakes*, 187-198.

- 778 Kloss, C., Goniva, C., Hager, A., Amberger, S., & Pirker, S. (2012). Models,  
779 algorithms and validation for opensource dem and cfd-dem. *Prog. Comput.*  
780 *Fluid Dyn.*, 12(2-3), 140-152.
- 781 Knuth, M., & Marone, C. (2007). Friction of sheared granular layers: Role of particle  
782 dimensionality, surface roughness, and material properties. *Geochem. Geophys.*  
783 *Geosyst.*, 8(3). doi:doi:10.1029/2006GC001327
- 784 Kozłowska, M., Brudzinski, M. R., Friberg, P., Skoumal, R. J., Baxter, N. D., &  
785 Currie, B. S. (2018). Maturity of nearby faults influences seismic hazard from  
786 hydraulic fracturing. *Proceedings of the National Academy of Sciences*, 115(8), E1720-  
787 E1729. doi:10.1073/pnas.1715284115
- 788 Kwiatek, G., Goebel, T. H. W., & Dresen, G. (2014). Seismic moment tensor and b  
789 value variations over successive seismic cycles in laboratory stick-slip  
790 experiments. *Geophys. Res. Lett.*, 41(16), 5838-5846.  
791 doi:doi:10.1002/2014GL060159
- 792 Leeman, J. R., Saffer, D. M., Scuderi, M. M., & Marone, C. (2016). Laboratory  
793 observations of slow earthquakes and the spectrum of tectonic fault slip  
794 modes. *Nat Commun*, 7, 11104. doi:10.1038/ncomms11104
- 795 <https://www.nature.com/articles/ncomms11104#supplementary-information>
- 796 Mair, K., Frye, K. M., & Marone, C. (2002). Influence of grain characteristics on the  
797 friction of granular shear zones. *J. Geophys. Res.*, 107(B10).  
798 doi:10.1029/2001jb000516
- 799 Makedonska, N., Sparks, D. W., Aharonov, E., & Goren, L. (2011). Friction versus  
800 dilation revisited: Insights from theoretical and numerical models. *J. Geophys.*  
801 *Res.*, 116(B9). doi:10.1029/2010jb008139
- 802 Marone, C. (1998a). The effect of loading rate on static friction and the rate of fault  
803 healing during the earthquake cycle. *Nature*, 391(6662), 69-72.  
804 doi:10.1038/34157
- 805 Marone, C. (1998b). Laboratory-derived friction laws and their application to seismic  
806 faulting. *Annu. Rev. Earth Planet. Sci.*, 26, 643-696.  
807 doi:10.1146/annurev.earth.26.1.643
- 808 Marone, C., Carpenter, B. M., & Schiffer, P. (2008). Transition from rolling to  
809 jamming in thin granular layers. *Physical Review Letters*, 101(24), 248001.

- 810 Marone, C., & Kilgore, B. (1993). Scaling of the critical slip distance for seismic  
811 faulting with shear strain in fault zones. *Nature*, 362, 618.  
812 doi:10.1038/362618a0
- 813 Marone, C., Raleigh, C. B., & Scholz, C. H. (1990). Frictional behavior and  
814 constitutive modeling of simulated fault gouge. *J. Geophys. Res.*, 95(B5), 7007-  
815 7025. doi:10.1029/JB095iB05p07007
- 816 McLaskey, G. C., & Glaser, S. D. (2011). Micromechanics of asperity rupture during  
817 laboratory stick slip experiments. *Geophys. Res. Lett.*, 38(12).  
818 doi:doi:10.1029/2011GL047507
- 819 Mead, W. J. (1925). The geologic role of dilatancy. *The Journal of Geology*, 33(7), 685-  
820 698.
- 821 MiDi, G. D. R. (2004). On dense granular flows. *Eur. Phys. J. E*, 14(4), 341-365.  
822 doi:10.1140/epje/i2003-10153-0
- 823 O'Sullivan, C. (2011). *Particulate discrete element modelling: A geomechanics perspective*. Spon  
824 Press.
- 825 Ohnaka, M. (2003). A constitutive scaling law and a unified comprehension for  
826 frictional slip failure, shear fracture of intact rock, and earthquake rupture. *J.*  
827 *Geophys. Res.*, 108(B2). doi:doi:10.1029/2000JB000123
- 828 Ohnaka, M., & Shen, L.-f. (1999). Scaling of the shear rupture process from  
829 nucleation to dynamic propagation: Implications of geometric irregularity of  
830 the rupturing surfaces. *J. Geophys. Res.*, 104(B1), 817-844.  
831 doi:doi:10.1029/1998JB900007
- 832 Park, J.-W., & Song, J.-J. (2009). Numerical simulation of a direct shear test on a rock  
833 joint using a bonded-particle model. *Int. J. Rock Mech. Min. Sci.*, 46(8), 1315-  
834 1328. doi:https://doi.org/10.1016/j.ijrmms.2009.03.007
- 835 Pierre, R., S., B. H., Romain, J., & Raúl, M. (2018). Fast and slow slip events emerge  
836 due to fault geometrical complexity. *Geophys. Res. Lett.*, 45(10), 4809-4819.  
837 doi:doi:10.1029/2018GL077579
- 838 Rabinowicz, E. (1956). Stick and slip. *Sci. Am.*, 194(5), 109-119.

- 839 Rathbun, A. P., & Marone, C. (2010). Effect of strain localization on frictional  
840 behavior of sheared granular materials. *J. Geophys. Res.*, *115*(B1).  
841 doi:10.1029/2009jb006466
- 842 Rathbun, A. P., Renard, F., & Abe, S. (2013). Numerical investigation of the interplay  
843 between wall geometry and friction in granular fault gouge. *J. Geophys. Res.*,  
844 *118*(3), 878-896. doi:doi:10.1002/jgrb.50106
- 845 Renard, F., & Candela, T. (2017). Scaling of fault roughness and implications for  
846 earthquake mechanics *Fault zone dynamic processes: Evolution of fault properties*  
847 *during seismic rupture, geophysical monograph 227*: John Wiley & Sons, Inc.
- 848 Renard, F., Voisin, C., Marsan, D., & Schmittbuhl, J. (2006). High resolution 3d laser  
849 scanner measurements of a strike-slip fault quantify its morphological  
850 anisotropy at all scales. *Geophys. Res. Lett.*, *33*(4).  
851 doi:doi:10.1029/2005GL025038
- 852 Rijsingen, E., Lallemand, S., Peyret, M., Arcay, D., Heuret, A., Funicello, F., & Corbi,  
853 F. (2018). How subduction interface roughness influences the occurrence of  
854 large interplate earthquakes. *Geochem. Geophys. Geosyst.*  
855 doi:doi:10.1029/2018GC007618
- 856 Rivière, J., Lv, Z., Johnson, P. A., & Marone, C. (2018). Evolution of b-value during  
857 the seismic cycle: Insights from laboratory experiments on simulated faults.  
858 *Earth Planet. Sci. Lett.*, *482*, 407-413.  
859 doi:https://doi.org/10.1016/j.epsl.2017.11.036
- 860 Rosenau, M., Corbi, F., & Dominguez, S. (2017). Analogue earthquakes and seismic  
861 cycles: Experimental modelling across timescales. *Solid Earth*, *8*(3), 597-635.  
862 doi:10.5194/se-8-597-2017
- 863 Rouet-Leduc, B., Hulbert, C., Lubbers, N., Barros, K., Humphreys, C. J., & Johnson,  
864 P. A. (2017). Machine learning predicts laboratory earthquakes. *Geophys. Res.*  
865 *Lett.*, *44*(18), 9276-9282. doi:10.1002/2017GL074677
- 866 Sagy, A., Brodsky, E. E., & Axen, G. J. (2007). Evolution of fault-surface roughness  
867 with slip. *Geology*, *35*(3), 283-286. doi:10.1130/G23235A.1
- 868 Schubnel, A., Nielsen, S., Taddeucci, J., Vinciguerra, S., & Rao, S. (2011). Photo-  
869 acoustic study of subshear and supershear ruptures in the laboratory. *Earth*  
870 *Planet. Sci. Lett.*, *308*(3), 424-432.  
871 doi:https://doi.org/10.1016/j.epsl.2011.06.013

- 872 Sheng, Y., Lawrence, C. J., Briscoe, B. J., & Thornton, C. (2004). Numerical studies  
873 of uniaxial powder compaction process by 3d dem. *Eng. Computation.*, 21(2-4),  
874 304-317. doi:10.1108/02644400410519802
- 875 Shimamoto, T. (1979). *Experimental studies of simulated gouge and their application to studies*  
876 *of natural fault zones*. Paper presented at the Conference VIII: Analysis of Actual  
877 Fault Zones in Bedrock, Natl. Earthquake Hazards Reduct. Program, Menlo  
878 Park, Calif.
- 879 Shojaee, Z., Brendel, L., Török, J., & Wolf, D. E. (2012). Shear flow of dense  
880 granular materials near smooth walls. Ii. Block formation and suppression of  
881 slip by rolling friction. *Phys. Rev. E*, 86(1), 011302.  
882 doi:10.1103/PhysRevE.86.011302
- 883 Tal, Y., & Hager, B. H. (2018). The slip behavior and source parameters for  
884 spontaneous slip events on rough faults subjected to slow tectonic loading. *J.*  
885 *Geophys. Res.*, 123(2), 1810-1823. doi:doi:10.1002/2017JB014737
- 886 Tal, Y., Hager, B. H., & Ampuero, J. P. (2018). The effects of fault roughness on the  
887 earthquake nucleation process. *J. Geophys. Res.*, 123(1), 437-456.  
888 doi:doi:10.1002/2017JB014746
- 889 Wang, K., & Bilek, S. L. (2014). Invited review paper: Fault creep caused by  
890 subduction of rough seafloor relief. *Tectonophysics*, 610, 1-24.  
891 doi:https://doi.org/10.1016/j.tecto.2013.11.024
- 892 Wang, W.-J., Kong, X.-Z., & Zhu, Z.-G. (2007). Friction and relative energy  
893 dissipation in sheared granular materials. *Phys. Rev. E*, 75(4), 041302.
- 894 Xia, K., Rosakis, A. J., & Kanamori, H. (2004). Laboratory earthquakes: The sub-  
895 rayleigh-to-supershear rupture transition. *Science*, 303(5665), 1859-1861.  
896 doi:10.1126/science.1094022
- 897 Zielke, O., Galis, M., & Mai, P. M. (2017). Fault roughness and strength  
898 heterogeneity control earthquake size and stress drop. *Geophys. Res. Lett.*, 44(2),  
899 777-783. doi:doi:10.1002/2016GL071700
- 900 Zoback, M., Hickman, S., & Ellsworth, W. (2010). Scientific drilling into the san  
901 andreas fault zone. *Eos, Transactions American Geophysical Union*, 91(22), 197-  
902 199.

903