PRESERVATION OF AUTOGENIC PROCESSES AND ALLOGENIC FORCINGS WITHIN SET-SCALE AEOLIAN ARCHITECTURE II: THE SCOUR-AND-FILL DOMINATED JURASSIC PAGE SANDSTONE, ARIZONA, USA

RUNNING TITLE: AUTOGENIC PROCESSES AND ALLOGENIC FORCINGS IN AEOLIAN STRATIGRAPHY II

BENJAMIN T. CARDENAS[1], GARY KOCUREK[1], DAVID MOHRIG[1], TRAVIS SWANSON[2], CORY M. HUGHES[1*], and SARAH C. BROTHERS[1**]

[1]Jackson School of Geosciences, University of Texas at Austin, Austin, TX, USA

[2]Department of Earth Science, Rice University, Houston, TX, USA

[*]now at Department of Geology, Western Washington University, Bellingham, WA, USA

[**]now at National Academies Space Studies Board, Washington, D.C., USA

Corresponding Author Email: benjamin.cardenas@utexas.edu

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ABSTRACT

The stratigraphic architecture of aeolian sandstones is thought to record signals originating from both autogenic dune behavior and allogenic environmental boundary conditions within which the dune field evolves. Mapping of outcrop-scale surfaces and sets of cross-strata between these surfaces for the Jurassic Page Sandstone near Page, Arizona, USA, demonstrates that the stratigraphic signature of autogenic behavior is captured by variable scour depths and subsequent fillings, whereas the dominant signatures of allogenic boundary conditions are associated with antecedent surface topography and variable water-table elevations. At the study area, the Page Sandstone is ~ 60 m thick and is separated from the underlying Navajo Sandstone by the J-2 regional unconformity with meters of relief. Thin, climbing sets of cross-strata of the basal Page representing early dune-field accumulations fill J-2 depressions. In contrast, the overlying lower and middle Page consist of cross-strata that are one to a few meters thick, and packaged between outcrop-scale bounding surfaces. These bounding surfaces have been previously correlated to high stand deposits of the adjacent Carmel sea and at this site possess meters of erosional relief produced by dune scour. Notably absent within packages of cross-strata bounded by these outcrop-scale surfaces are strata of early dune-field accumulations, any interdune deposits, and climbing-dune strata. Instead, these packages preserve a scour-and-fill architecture created by large migrating dunes migrating within a dry, mature, dune field undergoing negligible bed aggradation. Any record of early phases of dune-field construction for the lower and middle Page are interpreted to have been cannibalized by the deepest scours of later, large dunes. Interpretations are independently supported by the relatively large coefficients of variation in lower and middle Page set thicknesses, which are consistent with set production by successive deepest trough scours, and the relatively low coefficient of variation for the

depression-filling basal Page sets consistent with a significant component of bed aggradation. Numerical modeling presented here and more completely in the companion paper demonstrates how this cannibalization of early-phase stratigraphy is an expected outcome of autogenic dune-growth processes, and that early-phase strata can be preserved within antecedent depressions. Relative rise of the inland water table from basin subsidence and Carmel sea level forced preservation of multiple stacked packages composed of scour-and-fill architecture. Without these allogenic forcings, the Page would be little more than an erosional surface.

INTRODUCTION

Aeolian dune fields develop over time as a result of autogenic processes that occur within a set of environmental (allogenic) boundary conditions. Autogenic processes inherent to a field of migrating dunes include dune interactions (Werner 1995; Ewing and Kocurek 2010a; Kocurek et al. 2010; Gao et al. 2015a), dune deformation with migration (Pedersen et al. 2015; Swanson et al. 2016), and dune scour of the substrate (Paola and Borgman 1991; Bridge and Best 1997). Common allogenic boundary conditions for aeolian systems include the presence or absence of a near-surface water table (Crabaugh and Kocurek 1993; Kocurek and Havholm 1993), direction and magnitude of sediment-transporting winds (Rubin 1987; Rubin and Hunter 1987; Ping et al. 2014; Swanson et al. 2017), sediment availability (Courrech du Pont et al. 2014; Gao et al. 2015b), and geometry of the sediment source and basin shape (Ewing and Kocurek 2010b).

The general trend in dune-field development is for many, small, closely spaced dunes to coalesce into fewer, larger, widely spaced dunes over time and space through constructive dune interactions (Ewing and Kocurek 2010a; Eastwood et al. 2011; Gao et al. 2015a; Day and Kocurek, 2018). Aeolian strata record these interaction kinematics (Brothers et al. 2017; Day and

Kocurek 2017), and may potentially preserve strata associated with any phase of this development. Conversely, the presence or absence of preserved accumulations associated with different developmental stages are useful data for interpreting the allogenic and autogenic forces that led to its accumulation and preservation (Kocurek and Day 2017).

In this work, we examine the Jurassic Page Sandstone near Page, Arizona, USA, through the collection of detailed field mapping and topographic measurements. Facies interpretations, stratal architecture, and bounding surface topography are used to demonstrate that the Page sets of cross-strata are dominated by a scour-and-fill architecture constructed by relatively large, mature dunes. Moreover, this later-phase of dune construction cannibalized most strata that may have accumulated during earlier phases of dune-field development, as well as having scoured into underlying strata from previous constructional events. Even so, two examples of earlier dune-field phases and their distinct cross-strata facies were preserved in local, pre-existing topographic lows. Variable trough scour depth was the dominant autogenic control on preserved stratigraphy, whereas antecedent topography and depth to water table are demonstrated to have been the dominant allogenic controls on the architecture of the Page Sandstone. Interpretation was aided by a numerical model coupling dune morphodynamics and stratigraphy, which is the focus of the companion paper (Swanson et al. this issue).

Geologic Context and Previous Work

Jurassic aeolian formations of the Colorado Plateau are among the most studied aeolian sandstones in the world (Blakey et al. 1983, 1988; Rodríguez-López et al. 2014), and literature discussing the Page Sandstone is extensive. The Page Sandstone preserves a time series of NE-SW trending dune fields situated between the Monument Upwarp and the Carmel inland sea (Blakey et al. 1988; Riggs and Blakey 1993; Peterson 1994) (Fig. 1A) during the Middle

Jurassic, 171.5–169.5 Ma (Blakey and Parnell 1995; Dickinson et al. 2010). Facies within the Carmel Formation represent shallow marine, sabkha, and fluvial settings, and these intertongue with aeolian Page accumulations over a belt ~75 km wide running parallel to the paleocoastline (Havholm et al. 1993; Blakey et al. 1996; Taggart et al. 2010). Intertonguing of the Carmel coastal complex with the western portions of the Page Sandstone is interpreted to represent the interplay of tectonic subsidence within the Utah-Idaho trough, changes in sediment supply, and sea level (Blakey et al. 1996).

The Page Sandstone is separated from the underlying Navajo Sandstone by the J-2 surface, one of six regional unconformities formed across the greater Colorado Plateau during the Jurassic (Pipiringos and O'Sullivan 1978). The J-2 surface near Page, Arizona, is characterized by large polygonal fractures, diagenetic chert nodules, and meters of erosional relief (Pipiringos and O'Sullivan 1978; Kocurek and Hunter 1986; Swezey 1991; Kocurek et al. 1991). The Page is overlain by the Carmel Formation, representing eastward progradation of the Carmel fluvial and coastal complex (Blakey et al. 1996).

Previous work has produced hundreds of correlated vertical sections across the entirety of the Page Sandstone (Capps 1990; Jones 1990; Havholm 1991; Kocurek et al. 1991; King 1992; Havholm et al. 1993; Jones and Blakey 1993; Havholm and Kocurek 1994; Blakey et al. 1996). The Page has been informally divided into a basal, lower, middle, and upper unit by Havholm et al. (1993), and these divisions correlate with formal stratigraphic names used by Blakey et al. (1996) (Fig. 1B). The informal units were defined using formation-scale, erosional bounding surfaces. These surfaces are characterized by polygonal fractures, interpreted as having developed in evaporite-cemented sand, and/or overlying wavy bedding interpreted as sabkha deposits (Figs. 2, 3; Kocurek and Hunter 1986; Havholm et al. 1993; Havholm and Kocurek

1994). Both of these features have been used as stratigraphic proxies for the paleo-water table (i.e., “Stokes surfaces” of Fryberger et al. 1988; “super surfaces” of Kocurek 1988). Each of these surfaces can be traced westward to where it is overlain by a transgressive tongue of the Carmel (Havholm et al. 1993; Havholm and Kocurek 1994; Blakey et al. 1996) (Fig. 1B). Because the Carmel transgressive tongues represent relative high stands of the Carmel sea, and their correlative inland surfaces are marked by features associated with the water table, the surfaces are interpreted to represent the elevation of the coastal water table, which rose in response to the adjacent sea-level rise (Havholm and Kocurek 1994; Blakey et al. 1996; Kocurek et al. 2001). The surfaces themselves are interpreted as having been formed by deflation down to the water table during the high stands in sea level when sediment availability was limited. Conversely, Page dune systems are thought to have developed during low stands in sea level that provided greater sand availability. Each body of Page cross-strata bounded by interpreted deflationary surfaces is therefore inferred to represent aeolian strata accumulated during a low stand and preserved as a consequence of a rising continental water table that protected it from wind-blown deflation. The thickness of a preserved accumulation reflects the cumulative effects of subsidence and relative sea-level rise, as reflected in the continental water table (Havholm and Kocurek 1994; Blakey et al. 1996).

Aeolian Stratification Types

The three basic stratification types that compose aeolian cross-strata are grainfall, grainflow, and wind-ripple deposits (Hunter, 1977). We used the architecture of these stratification types to interpret relative dune size from preserved sets of cross-strata. Grainfall deposits form when saltating grains are directly deposited on the lee faces of dunes. Grainfall is typically greatest immediately downwind of the brink and the deposition rate decreases down the lee face. As a result, the lee slope steepens until the deposit builds to the angle of initial yield, triggering an avalanche, or grainflow (Allen 1970a; Hunter 1977; Kocurek and Dott 1981; McDonald and Anderson, 1995; Nield et al. 2017). Grainflow deposits are primarily located along portions of the lee face aligned close to perpendicular to the sediment-transporting wind direction (Eastwood et al. 2012). Wind-ripple laminations form via wind-ripple migration and aggradation along dune lee faces (Hunter 1977), particularly those segments obliquely oriented to the regional wind direction (Eastwood et al. 2012).

Commonly, aeolian cross-strata are composed of alternating packages of grainflow deposits and wind-ripple laminations, representing either seasonal changes in wind direction and/or changes in dune shape and crest orientation (Hunter and Rubin 1983). Because aeolian sets of cross-strata typically represent only the basal portions of dune lee faces, grainfall deposits are less frequently preserved in sets representing large dunes. Grainfall deposits are, however, common in cross-strata representing small dunes, where the grainfall apron extended to the base of the dunes (Kocurek and Dott 1981). The Page Sandstone exhibits all three stratification types, and their typical arrangements within the range of set thicknesses are presented as Figures 4A to C and 5A to B.

Aeolian set architecture

The architecture of aeolian sets of cross-strata between outcrop-scale bounding surfaces preserve a record of dune-field kinematics (Fig. 6A; Allen 1970b; Rubin and Hunter 1982). In cases of dune climb, where trains of migrating dunes must climb over the accumulations of downwind dunes, set boundaries are expected to originate from a lower, outcrop-scale bounding surface and rise at some measurable angle. In cases where sets are filling a pre-existing topographic low, the architecture is defined by set boundaries that originate at an upper, outcrop-scale bounding surface and descend, ultimately downlapping onto a lower, outcrop-scale bounding surface that hosts the antecedent topographic low. Finally, in cases where variable depth of trough scour creates space that is subsequently filled by smaller dunes in the train, sets boundaries persist for limited distances before being cut out by other set boundaries rather than by outcrop-scale bounding surfaces.

The architecture of set boundaries is not the only data type used here to reconstruct the filling history. Distributions on set thickness also contain information as to how the stratification formed (Paola and Borgman 1991; Bridge and Best 1997; Jerolmack and Mohrig, 2005). In particular, distributions of cross-strata produced by the scour-and-fill process (Fig. 6A) record an asymptotically high value for the coefficient of variation of bed thicknesses, while sets produced during net bed aggradation (climbing and downlapping cases in Fig. 6A) have distinctly smaller values for coefficient of variation associated with them. Here we have combined set architecture with set statistics to increase our confidence in the reconstructed processes that produced the observed stratigraphy.

METHODS

Field Work

Three outcrops in and near Page, Arizona, were selected based upon their accessibility and orientation relative to the paleo-transport direction determined from a measured distribution of cross-strata dip-directions (Fig. 7). The Ferry Swale outcrop was split into two sections; one segment oriented perpendicular to the general dune migration direction and the other oblique to the dune migration direction. The Manson Mesa and Golf Course sections were each oriented roughly parallel to the transport direction. Cross-strata set boundaries and outcrop-scale bounding surfaces were surveyed using a total station at each outcrop location. A GPS point was taken at each base station location so that all surveyed points could be converted to UTM coordinates. Each outcrop-scale erosional surface was then correlated with vertical sections from Havholm (1991). In addition to mapping surfaces, 90 measurements of foreset dip direction were taken using a Brunton compass and averaged with the Circular Statistics Toolbox in MATLAB (Berens 2012). Thicknesses of 161 cycles of wind-ripple and stacked grainflow strata were collected from cross-strata. All of these measurements were collected from the basal Page (PB), lower Page (PL), and middle middle Page (PMm) intervals of the Page Sandstone (Fig. 1). For simplicity, we will refer to the middle middle Page as the middle Page. These divisions make up the strata bounded by the underlying J-2 and overlying S-rm surfaces (Fig. 1B). The upper Page, unit PUI of Figure 1, is composed of sets of compound-dune cross-strata that are distinctively different from those of the underlying units and are not part of this study.

Data Processing

Combining field maps, with GPS datums and annotated total-station points, a digital GIS project including all of the outcrop locations was constructed. Continuous surfaces were interpolated from point data defining each outcrop-scale bounding surface and cross-strata set boundary using kreiging method that completely preserves the XYZ input data. The spatial resolution for resulting digital elevation models (DEMs) at Manson Mesa, the Golf Course, and Ferry Swale is 0.55 m, 0.65 m, and 2 m.

Using the GIS, vertical stratigraphic sections were constructed at 20 m intervals across each outcrop (Figs. 8- 9 and S1-S3). The 3-D mapping data allowed surfaces within each cross-section to be accurately correlated. Measurements of set thickness and outcrop-scale bounding-surface relief were made from the generated cross-sections. Set geometry and its relationship to the nearest bounding surface was also categorized from the cross-sections (Fig. 6A).

Numerical Modeling

Swanson et al. (this issue) numerically model surface topography, dune migration, dune interactions, and cross-strata accumulation by coupling bed topography, bed shear stress, and sediment transport. Using this model, several outcrop-scale bounding surfaces and stratigraphic packages have been re-created in a 2-D panel. Observations of this synthetic stratigraphy are compared to outcrop interpretations (Fig. 10; model parameters in Table S1.) To create the synthetic Page stratigraphy, a dune field and its accumulations are developed for a set period of time during which an initial rough surface develops into small, early-phase dunes, which develop into larger, more mature dunes. Following that time, a relatively flat outcrop-scale bounding surface is formed through the stratigraphy that has developed. Then, another dune field and its accumulations develop on top of this outcrop-scale bounding surface. Four episodes of aeolian accumulation are modeled in this scenario, creating four outcrop-scale packages bound by three outcrop-scale bounding surfaces (Fig. 10). The synthetic stratigraphy is color-coded by age relative to each accumulation episode.

RESULTS

Outcrop-Scale Architecture

At the three outcrops (Figs. 8-9 and S1-S3) the Page Sandstone is composed of packages of aeolian cross-strata partitioned by the formation-scale surfaces of Havholm et al. (1993). Within the studied interval spanning from the J-2 to the S-rm surfaces (Fig. 1), the Page is nowhere thicker than 60 m and is largely composed of meter-scale beds of cross-strata in the lower and middle Page (Fig. 4A), with local preservation of thinner sets in the basal Page (Figs. 5A and 11A-D). In addition to the formation-scale surfaces that define the informal Page units, Havholm et al. (1993) identified less-continuous surfaces that are typically truncated laterally. For this study, the term "outcrop-scale bounding surface" refers to both the formation-scale and less-continuous super surfaces identified by Havholm et al. (1993).

Outcrop-Scale Bounding Surfaces

Figures 8-9 and S1-S3 present the structure of the outcrop-scale bounding surfaces surveyed at the three outcrop locations (Fig. 2). The J-2 surface is characterized by wedge-shaped fractures into the underlying Navajo Formation, pebble-sized, diagenetic chert nodules replacing evaporites within the uppermost Navajo strata, and up to 10 m of local erosional relief (e.g., Ferry Swale; Fig. S1). The outcrop-scale bounding surfaces within the lower and middle Page also preserve erosional relief that varies between a minimum of 1.2 m to as much as 13.5 m (Fig. 12A). On both walls of the Ferry Swale outcrop (Figs. 8-9), this relief has a scalloped geometry. Horizontal distances from the adjacent high points on either side of a scallop were measured on the western, perpendicular-to-transport wall (Fig. 8). The widths of sets filling these scallops were measured on the same wall as the horizontal distance between pinchouts. Scallop widths (n = 23) have a mean value of 72 m and a standard deviation of 34 m. Sets filling these scallops (n = 33) have a similar mean width of 64 m and a greater standard deviation of 81 m.

Outcrop-scale bounding surfaces are commonly associated with two unique deposit types. Wavy-laminated sandstones with variable thicknesses ranging up to 0.4 m occur discontinuously along outcrop-scale bounding surfaces in the lower and middle Page (Fig. 2A-D). In the basal Page, the wavy laminated sandstones were identified both overlying the J-2 surface and in between aeolian cross-sets, and these strata are generally thicker and more continuous than in the higher Page units. As a whole, the red, wavy-laminated strata have been interpreted as sabkha deposits that formed along surfaces deflated to the near-surface water table (Havholm et al. 1993; Havholm and Kocurek 1994). The second deposit type primarily associated with outcrop-scale bounding surfaces are prominent, wedge-shaped fracture fills. In exposures that provide cross-sectional views, the fractures cross-cut sets of cross-stratified sandstone (Fig. 3A-C). These fractures narrow downward, although not always in a linear fashion. In plan-view, the fracture fills take a more polygonal shape (Fig. 3A). Sandstone fill within the fractures is vertically-laminated or structureless (Fig. 3B-C). The tops of the wedges terminate against outcrop-scale bounding surfaces. These sandstone wedges have been interpreted as sand-filled, salt-cemented thermal contraction polygonal fractures (Kocurek and Hunter 1986).

Cross-Set Architecture

The outcrop-scale bounding surfaces truncate and bound packages of cross-stratified beds, referred to hereafter as "outcrop-scale packages." Within these packages, individual sets of cross-strata are defined by upper and lower set boundaries that are less continuous than the outcrop-scale surfaces. The architecture of outcrop-scale packages varies from the basal to middle Page.

Basal Page Architecture.---The thin sets of cross-stratified sandstone (Fig. 5A-B) are relatively uncommon and limited to the basal Page at the outcrops studied. Set thicknesses range from 0.03 to 0.45 m with a mean of 0.16 m and a standard deviation of 0.10 m (Fig. 5C; n = 80). Individual grainflow deposits are well defined by their characteristic blade shape and are separated by mm-thick grainfall deposits (Fig. 5B). Alternating grainflow and grainfall deposits define the cross-stratification in these sets. The mean grainflow deposit thickness is 18 mm with a standard deviation of 9 mm (Fig. 5D; n = 36).

In contrast to the lower and middle Page, the sets of the basal Page are stacked tens of sets high to form packages several meters thick. This architecture occurs at Ferry Swale and north of the Golf Course outcrop (Fig. 11A-D and 13). At the Ferry Swale outcrop, these thin sets are located filling a local J-2 depression (Fig. S1 shows full J-2 topography at Ferry Swale). North of the Golf Course outcrop, basal Page sets probably also fill a depression on the J-2 surface, but the contact with the Navajo is not exposed. At both locations, these packages of stacked, thin sets are laterally scoured and filled by thicker sets of the lower Page (Fig. 11A-D and 13).

The stacked, thin sets of the basal Page are the focus of our comparison between the basal Page and the lower and middle Page. However, parts of the basal Page also consist of alternating sabkha and cross-sets, as reported by Havholm et al. (1993) but not analyzed here.

Lower and Middle Page Architecture.---The predominant facies composing the lower and middle Page at all three sites (Fig. 7) are thick sets of cross-stratified sandstones (Fig. 4A) composed of foresets consisting of stacked grainflow strata separated by intervals of wind-ripple laminae (Fig. 4B). Where grainflow cross-strata do not extend to the base of the set, these yield downward to low-angle wind-ripple strata that form toesets (Fig. 4C). Individual grainflow deposits cannot be distinguished, but rather form amalgamations of multiple grainflow deposits. From 461 measurements made using the cross-sections (Figs. 8-9 and S1-S3), these sets of cross-strata are on average 2.3 m thick, and occasionally exceed 10 m (Fig. 4D). The standard deviation of set thicknesses is 2.1 m. The thickness of individual sets is laterally variable at the outcrop scale (Fig. 8-9), and sets are typically truncated by adjacent sets of cross-strata, limiting lateral continuity (Fig. 14A-B). The thicknesses of grainflow deposit packages were measured from several sets of these cross-strata (e.g., Fig. 4B), and range from 0.01 to 1.30 m with an average value and standard deviation of 0.19 m (Fig. 5D). Ninety measurements of foreset dip direction were taken from this facies and serve as a proxy measurement of paleo-transport direction. The mean of these measurements is 143°.

The number of sets stacked between outcrop-scale surfaces at any vertical section of the lower and middle Page ranges from 1 to 5 (Figs. 8-9 and S1-S3). There is not a strong correlation between the number of stacked sets and the thickness of the associated outcrop-scale package (Fig. 12B). There is, however, a near one-to-one correlation between the thickness of an outcrop-scale package and the maximum amount of relief along the associated lower outcrop-scale bounding surface (Fig. 12A).

Using the cross-sections created for each outcrop (Figs. 8-9 and S1-S3), each of the sets of the lower and middle Page was placed into one of the architectural categories defined in Figure 6A. The categories describe in-transport set architectures, so sets in the perpendicular-to-transport wall of Ferry Swale are not addressed here. The oblique-to-transport wall of Ferry Swale is assumed to exhibit these characteristics well enough that those sets are included in the count. Discontinuous, truncated sets fitting the scour-and-fill architecture occur most frequently, although all architectural categories are identified in outcrop (Fig. 6B).

Numerical Modeling Results

Several observations of the numerical model are relevant for interpreting the Page strata (Fig. 10). Firstly, the relief along the modeled outcrop-scale bounding surfaces develops coevally via dune scouring with the development of the dune field during each phase of aeolian sedimentation. Secondly, this relief is filled with the accumulations of the relatively late-phase dunes associated with the creation of the scour relief. Thirdly, except for antecedent topographic lows sequestering accumulations beneath the depths of scours to follow, there is almost no preservation of strata representing the early-phases of the dune fields (Fig. 10).

DISCUSSION

Cross-Strata Interpretation

The thin sets of cross-strata of the basal Page are interpreted to represent significantly smaller dunes than those that gave rise to the sets of the lower and middle Page. The thin (mean = 0.16 m, n = 80) sets of the basal Page are characterized by thin grainflow strata (Fig. 5D; mean = 17.6 mm, n = 37) separated by grainfall deposits (Fig. 5B). Although the relationship between grainflow deposit thickness and dune height may be modified by factors, especially wind speed (McDonald and Anderson 1995; Nickling et al. 2002; Nield et al. 2017; Cornwall et al. 2018), these consistently thin grainflow strata are characteristic of small dunes. Moreover, the common presence of grainfall deposits between grainflow strata indicates small dunes where grainfall deposits of observable thickness extended to the base of the set (Hunter 1977; Kocurek and Dott 1981).

Several observations indicate the lower and middle Page sets are the deposits of much larger dunes. Firstly, these sets on average are an order of magnitude thicker (mean = 2.3 m, n = 461) than those of the basal Page, and trough widths measure on average 64 m (n = 33). Secondly, the absence of grainfall deposits in the lower and middle Page sets is suggestive of larger dunes where significant grainfall seldom reaches basal lee faces (Hunter 1977; Kocurek and Dott 1981). Thirdly, the sets of the lower and middle Page show alternating grainflow and wind-ripple strata, as described in detail in Kocurek et al. (1991). This alternation of stratification types is commonly interpreted as seasonal cycles reflecting varying wind directions (Hunter and Rubin 1983). Large dunes formed with an abundant sand supply are oriented to the gross bedform-normal transport direction (Rubin and Hunter 1987), such that not all seasonal winds are transverse to the long-term crestline orientation. In contrast, repetitive packaging characteristic of multiple transporting wind directions are absent in the sets of the basal Page, which are interpreted to represent dunes small enough to reorient their crests with seasonal changes in the wind regime (e.g., bedform response time is shorter than wind direction cycles, Rubin and Hunter 1987; Ewing et al., 2015).

Dune fields begin as collections of protodunes that interact during migration and grow into larger dunes (e.g., Werner 1995; Ewing and Kocurek 2010a; Swanson et al. 2017). Therefore, the relatively small dunes associated with the deposition of the basal Page sets are interpreted to represent an earlier phase of dune-field development when compared to the dune fields that produced the larger sets of the lower and middle Page. Significantly, there are no preserved strata from a comparably early phase of dune-field development within the lower and middle Page. This is consistent with the model results (Fig. 10).

Stratigraphic Architecture and Outcrop-Scale Bounding Surface Topography

Architecture Preserved Within Antecedent J-2 Topography.---The J-2 regional surface represents regional erosion (Pipiringos and O'Sullivan 1978), and relief on this surface in the Page area is therefore not interpreted to be a product of dune scour associated with the migration of Page dunes. Rather, J-2 topography is interpreted as antecedent to the development of the Page system. Depressions along this surface had a measurable effect upon the architecture of the Page Sandstone by providing local accommodation space for sabkha deposits, packages of wind-ripple strata (Kocurek et al. 1991; Havholm et al. 1993), and local shallow pond deposits

(Swezey 1991). Antecedent depressions in the J-2 are also filled by stacks of thin cross-sets interpreted as representing small dunes of a juvenile dune field within the Page. These deposits aggraded here owing to local wind deceleration over the depression.

The J-2 antecedent relief also acted to prevent reworking of the early-phase strata by later, deep-scouring large dunes (Fig. 11A-D), as predicted by the numerical model (Fig. 10). Even so, packages of thin sets were partially scoured during reactivations of the Page dune fields and subsequently filled by thicker beds of cross-strata composed of grainflow and wind-ripple deposits. In the example at Ferry Swale, lateral scour of a section of basal Page (Fig. 11 C-D) is part of a large scoured depression that was subsequently filled by strata from at least two sides (Fig. 13). Near the Golf Course outcrop, several meters of stacked, thin sets of the basal Page are laterally truncated by a thick set of lower Page (Fig. 11 A-B). Although this section of basal Page was probably also formed within a J-2 depression, the J-2 surface is locally covered along adjacent outcrops of Navajo Sandstone. Aggradation of early phases of Page dune fields may have occurred beyond localized J-2 depressions but, ultimately, as dunes grew in size, these accumulations were likely completely cannibalized.

Scour-and-Fill Architecture and Coeval Development of Outcrop-Scale Bounding Surface Relief.---Several lines of evidence support a scour-and-fill type architecture for the lower and middle Page. In the perpendicular-to-transport wall of the Ferry Swale outcrop (Fig. 8), the similarity in width of the topographic scallops and the sets of cross-strata that fill the scallops indicates that (1) the topographic scallops were created by dune migration, and (2) the sets filling the scallops were deposited by the dunes that created the scour. This scour-and-fill architecture is in sharp contrast to that of the antecedent J-2 surface where the fill does not scale with the erosional relief. Because the outcrop-scale bounding surfaces represent a lateral continuum of distinct scallops (e.g., Fig. 8), these surfaces can be viewed as compound erosional surfaces. Although seen in outcrop as continuous surfaces with erosional relief, the outcrop-scale surfaces are interpreted to have formed with the passage of multiple dunes with varying scour depth. Moreover, consistent with the numerical model (Fig. 10), the ultimate relief on the surface would reflect the passage of the largest dunes with the deepest scours. This interpretation is consistent with the local truncation of sections of basal Page by the passage of large scours (Figs. 11, 13), and the lateral truncation of outcrop-scale packages by subsequent scour (Fig. 9). At the regional scale, Havholm et al. (1993) show the scour of the entire upper middle Page (PMu) along the regional traverse (Fig. 1B), and deep scours have been described elsewhere for the Page (Blakey 1988). For the latter examples, scour associated with the passage of trains of large dunes is not unreasonable, but other deflationary processes cannot be discounted.

Figure 12B compares the number of sets within an outcrop-scale package at a vertical section to the thickness of the outcrop-scale package at that same vertical section. There is not a significant correlation between the two measurements. A plot comparing the maximum relief across an outcrop-scale bounding surface to the thickness of the outcrop-scale package above that surface, however, shows a strong one-to-one relationship (Fig. 12A). The two plots demonstrate that although topographic depressions along outcrop-scale bounding surfaces do aid in the preservation of more of the rock record (i.e., a thicker outcrop-scale package), this thicker section does not correlate with a greater number of sets within an outcrop-scale package. Consistent with field observations, there is no selective preservation of more sets or of early-phase strata within deeper scours. This observation is again in contrast with the J-2 surface, where negative relief of the same scale houses accumulations of early-phase strata.

As also evident from Figure 12B, most outcrop-scale packages consist of only one or two sets of cross-strata. In the case of one set comprising the entire outcrop-scale package, the preserved record consists of the scoured trough and cross-strata fill of only the largest dune to migrate across the area. As predicted by the numerical model (Fig. 10), additional sets within an outcrop-scale package record the scour and fill of successive dunes that did not remove underlying sets. Consistent with the model, the dominant architecture of Page sets of cross-strata is one that represents scour-and-fill (Fig. 6).

Set thickness and set architecture within a preserved accumulation (i.e., an outcrop-scale package) are a function of climb angle and variability in scour depth (Fig. 12C; Paola and Borgman 1991). The dominance of scour-and-fill set architecture throughout the studied interval of the Page Sandstone suggests very low climb angles such that the deepest trough-scouring depths are the predominant control on set thickness (Fig. 12C; Paola and Borgman 1991; Bridge and Best 1997; Swanson et al. companion). This interpretation can be tested using relationships first established in Paola and Borgman (1991) and further generalized by Bridge and Best (1997) and Leclair et al. (1997). Mean set thickness, s_m, is a function of climb angle (δ), dune spacing (l), mean dune height (h_m), and the standard deviation of dune heights (h_{sd}^2):

$$s_m = l \tan(\delta) + 0.8225\ (h_{sd}^2/h_m). \qquad (1)$$

In Equation 1, $l \tan(\delta)$ describes the portion of mean set thickness due to climb (Allen, 1970b; Rubin and Hunter, 1982) and its associated net bed aggradation. The right-hand term, $0.8225(h_{sd}^2/h_m)$, describes the portion of mean set thickness produced by variability in dune size and thus, variability in scour depth. In the case of zero aggradation $(\delta = 0)$, s_m is only a function of variability in dune size:

$$s_m = 0.8225\ (h_{sd}^2/h_m). \qquad (2)$$

Bridge and Best (1997) note that with a climb angle of zero and gamma-distributed bedform heights, the standard deviation of set thicknesses (s_σ) is:

$$s_\sigma = 0.725\ (h_{sd}^2/h_m). \qquad (3)$$

Equations 2 and 3 can both be re-arranged to solve for h_{sd}^2/h_m, yielding

$$s_m/0.8225 = s_\sigma/0.725 \qquad (4)$$

for the case of no dune climb. Rearranging the terms in this equation yields a predicted value for the coefficient of variation of set thicknesses in the end-member case of zero net bed aggradation:

$$s_\sigma / s_m = 0.88 \qquad (5)$$

Because we have connected the dominance of scour-and-fill style sedimentation with a lack of significant aggradation using other arguments, the hypothesis is that Equation 5 should hold true for the lower and middle Page Sandstone assuming Page trough elevations were gamma distributed. It should be noted that Equation 5 is particularly useful in analyses of paleo-systems because it only requires measurements of set thicknesses that can be directly acquired in the field.

Using 461 measurements of thickness for sets of cross-strata preserved in the lower and middle Page yield an s_m = 2.34 m and s_σ = 2.10 m (Fig. 4D). Their ratio is 0.90, which is very close to the predicted 0.88 value of Equation 5. Importantly, the measured value is slightly greater than the predicted value. The inclusion of climb acts to reduce the value of the ratio until, with significant climb, the ratio is equal to that for the distribution of formative dune heights (Jerolmack and Mohrig 2005). This is consistent with the relatively long tail of the distribution, also observed by Jerolmack and Mohrig (2005) (Fig. 4D). In contrast, the stacked sets of the basal Page have an s_m = 0.16 m and s_σ = 0.10 m, yielding a coefficient of variation of 0.64. This value represents a higher aggradation rate and climb angle than is recorded in the lower and middle Page, consistent with the shorter tail of the set thickness distribution (Fig. 9C; Jerolmack and Mohrig 2005). This result from analysis of Page set thicknesses supports the argument for construction of these basal Page sets via the filling of antecedent J-2 depressions, while scour-and-fill dominated sets of the lower and middle Page are consistent with production under conditions of no climb or net bed aggradation. Set geometries and set statistics yield independent yet similar conclusions as to the mix of processes preserving cross-strata.

To demonstrate the universality of set thickness statistics in analyses of aeolian strata, the method was applied to set thickness measurements from the nearby Jurassic Entrada Sandstone, a type-example of an aggradational dune field with measurable positive climb angles (Carr-Crabaugh and Kocurek 1998; Kocurek and Day 2017). Thicknesses of 37 Entrada sets were taken from two vertical sections from Figure 2 of Kocurek and Day (2017), one characterized by downlapping sets filling a depression, and the other by climbing sets. For the Entrada, s_m = 2.10 meters, s_σ = 0.96 m, and their ratio is 0.46. This ratio value suggests sufficient aggradation such that the coefficient of variation for set thickness more closely reflects the coefficient of variation for the distribution of dune heights (Jerolmack and Mohrig 2005) rather than Equation 5. Indeed, Kocurek and Day (2017) interpret the Entrada strata as the product of allogenically-forced bed aggradation.

Role of Autogenic vs. Allogenic Controls on the Page Preserved Record

Prior work has addressed the difference in accumulation vs. preservation in the Page Sandstone at the regional scale by tracing formation-scale packages of sets of cross-strata and their bounding surfaces (Havholm and Kocurek 1994; Blakey et al. 1996). Our new work focuses upon the architecture of the sets within these formation-scale packages. Our measurements and interpretations are entirely consistent with those of the prior work, and additional observations informed by the results of Swanson et al. (this issue) allow for the reframing of the Page as a record of competing autogenic and allogenic signals, as is commonly done in fluvial and marine clastic depositional systems (Hajek and Straub 2017).

Autogenic variability in dune scour depth creates the scour-and-fill architecture observed throughout most of the Page Sandstone, identified here by set geometry and high variability in set thickness. There is no indication that any allogenic forcing beyond those allowing for transport have affected the architecture of the accumulations, indicating the water table was well below any point along the surface. This interpretation is consistent with previous work (Havholm and Kocurek 1994; Blakey et al., 1996), which interpreted Page dune constructional periods as occurring during low stands when the continental water table fell and sediment supply and availability increased. This new study argues that the periods of dune construction were characterized by scour of the substrate by migrating dunes with an accumulation that represents fill of these scours. Because the scours cannibalize underlying sediment, earlier stages of dune-field construction were reworked, as predicted by the numerical model (Fig. 10).

Autogenic processes of dune scour-and-fill alone, however, will not result in the preservation of stacked multiple dune constructional periods that represent the Page Sandstone. Indeed, without any additional forcing, the Page dune fields would be represented by an erosional surface and strata preserved just within depressions along the J-2 surface. In agreement with previous work (Havholm and Kocurek 1994; Blakey et al. 1996), preservation of Page accumulations is best accredited to allogenic forcing, which consisted of an episodic, but net progressive rise of the continental water table as a function of sea level in the adjacent Carmel.

The preserved accumulations capture those portions of the scour fill that were incorporated within the rising water table. Field evidence argues that high stands of the Carmel sea were characterized inland not only by a limit to deflation of the aeolian accumulations, but also a diminished sand supply such that dune fields were replaced by extensive polygonally fractured surfaces and sabkha deposition. Subsequent aeolian constructional periods during falling sea level and low stands were characterized by an influx of additional sand but also by scour into the surface, producing the compound erosional outcrop-scale surfaces. Scour was likely at least locally limited by the evaporite-cemented, polygonally-fractured surface (Havholm and Kocurek 1994). The set architecture of the preserved accumulations argue that scour-and-fill processes strongly dominated over dune climb.

The corollary to this interpretation is that antecedent topography and autogenically-formed dune scour topography along outcrop-scale bounding surfaces can be distinguished from one another. Relief autogenically formed from variability in scour depths typically cannibalizes early-phase accumulations, and scour widths scale well with the widths of the filling sets. The scour of the basal surface coeval with dune migration is consistent with the scour-and-fill architecture of the cross-sets themselves, as well as with prior interpretations regarding the formation of erosional outcrop-scale bounding surface relief (Havholm and Kocurek 1994). Broadly speaking, the preservation any non-scour-and-fill type architecture represents an allogenic forcing overcoming the autogenic, signal-shredding tendencies of the ancient Page dune fields. This is consistent with much of the current theory regarding the recording of environmental signals in stratigraphy (Jerolmack and Paola 2010; Paola et al. 2018), and framing future observations of aeolian stratigraphy in this manner will be informative of ancient environmental conditions on Earth and other planetary bodies with aeolian stratigraphic records, particularly Mars (e.g., Grotzinger et al. 2005; Milliken et al. 2014; Brothers et al. 2018; Banham et al. 2018; Day and Catling 2018; Anderson et al. 2018).

CONCLUSIONS

The Page Sandstone preserves a record of allogenic boundary conditions episodically overcoming an autogenic tendency of the dune field to shred its prior accumulations. Using field data and numerical modeling, the architecture of the Page Sandstone is characterized and linked to autogenic and allogenic forcings. Antecedent depressions along the J-2 surface allowed for the accumulation of early dune field deposits, identifiable by thin sets of cross-strata characterized by thin grainflow strata separated by basal grainfall deposits. These depressions also acted to locally preserve this early-phase stratigraphy from later scour. In contrast, periods of dune accumulation represented by the lower and middle Page were characterized by a lack of antecedent topography, a water table well below the depositional surface, and significant variation in dune scour depth. The compound erosional outcrop-scale bounding surfaces and scour-and-fill architecture of the accumulations are consistent with the reworking of any early dune field strata and argue that dune climb was subordinate to scour-depth variation in creating the set architecture. Topography developed by dune scour may be differentiated from antecedent relief by having similar scales of scour and fill. Numerical modeling presented here and in the companion paper (Swanson et al. this issue) highlight the scour of strata formed during early stages of dune field construction by subsequent autogenic dune scour. In contrast to the dominance of autogenic processes in shaping the architecture of the dune accumulations, their preservation requires allogenic forcing, which was the episodic and net progressive rise of the water table high associated with Carmel transgressions. Without episodic high stands driving preservation, the Page dune fields would likely be represented an amalgamated erosional surface and local preservation within J-2 depressions.

The methods used here to quantify the set architecture of the Page Sandstone are universal, and can be applied to any other aeolian sandstone where there is interest in understanding the formative environmental conditions of the associated ancient dune field. In addition to other Earth outcrops, aeolian sandstones on Mars could be characterized by similar methods using rover and remote sensing data.

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FIGURES

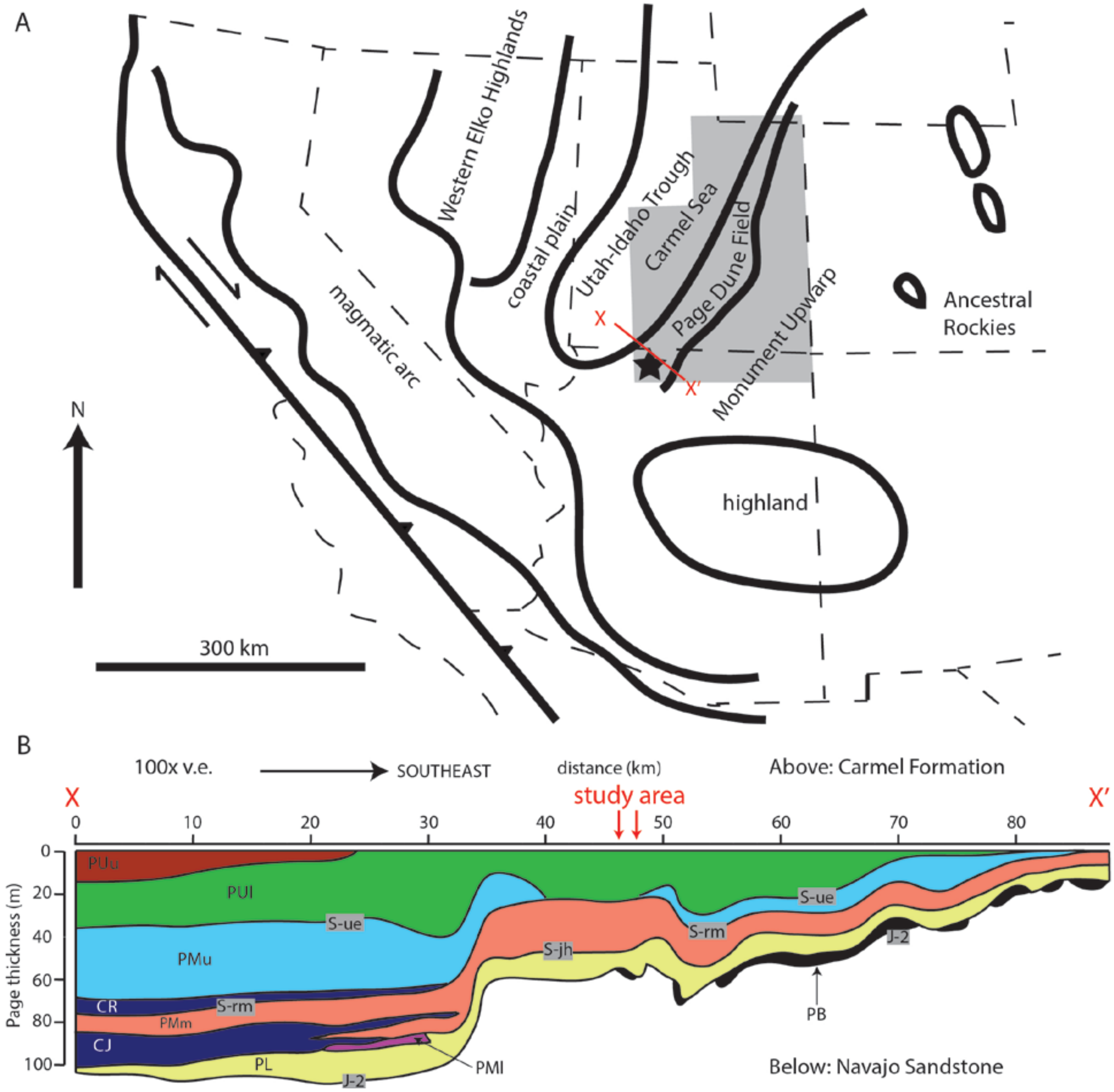

Figure 1 (18.2 cm x 17.7 cm) – A: Paleogeography of the Western Interior during the Middle Jurassic. The Page dune fields extended along the Carmel coastline, and were not present or were not preserved eastward over the Monument Upwarp. The gray box indicates the extent of Page and Carmel mapped by Havholm et al. (1993), Jones and Blakey (1993) and Blakey et al. (1996). The black star marks the location of Page, Arizona, and the outcrops discussed in this study. The red line X-X' shows the location of the cross-section in panel B. Modified after Havholm et al. (1993) and Peterson (1994). B: NW-SE cross-section X-X' showing the architecture of the Page Sandstone over more than 80 km. Figure is generalized after (Havholm et al., 1993, their Fig. 4D), but the cross-section has been hung from the Page/Carmel contact instead of the S-rm surface as in Havholm et al. (1993). Red arrows define the area of this study along the formation-scale cross-section. Overall, the cross-section shows the thinning of the Carmel/Page stratigraphic unit from the basin onto the flanks of the Monument Upwarp. Major formation-scale bounding surfaces separating informal units used by Havholm et al. (1993) are given. Page basal deposits (PB) and lower Page (PL) lie on the J-2 surface and are upward bounded by the S-jh surface. Westward the S-jh surface is overlain by a lower tongue of the Judd Hollow Member (CJ) of the Carmel Formation. Combined, the basal deposits and lower Page are the Harris Wash Member of Blakey et al. (1996). The middle Page (PM) or the Thousand Pockets Member of Blakey et al. (1996) is divided into lower (PMl), middle (PMm) and upper (PMu). PMl is bounded by a surface that is overlain by a higher, lesser tongue of the Judd Hollow. PMm is bounded by the S-rm surface, which is overlain westward by an upper tongue of the Judd Hollow (CR). PMu is bounded by the S-ue surface. Not shown in this cross-section, elsewhere westward the S-ue surface is overlain by the Crystal Creek Member of the Carmel. The upper Page (PUl and PUu, or the leche-e member of Blakey et al., 1996) consists of compound cross-strata not addressed in this study. The upper Page was completely transgressed and is overlain by younger members of the Carmel.

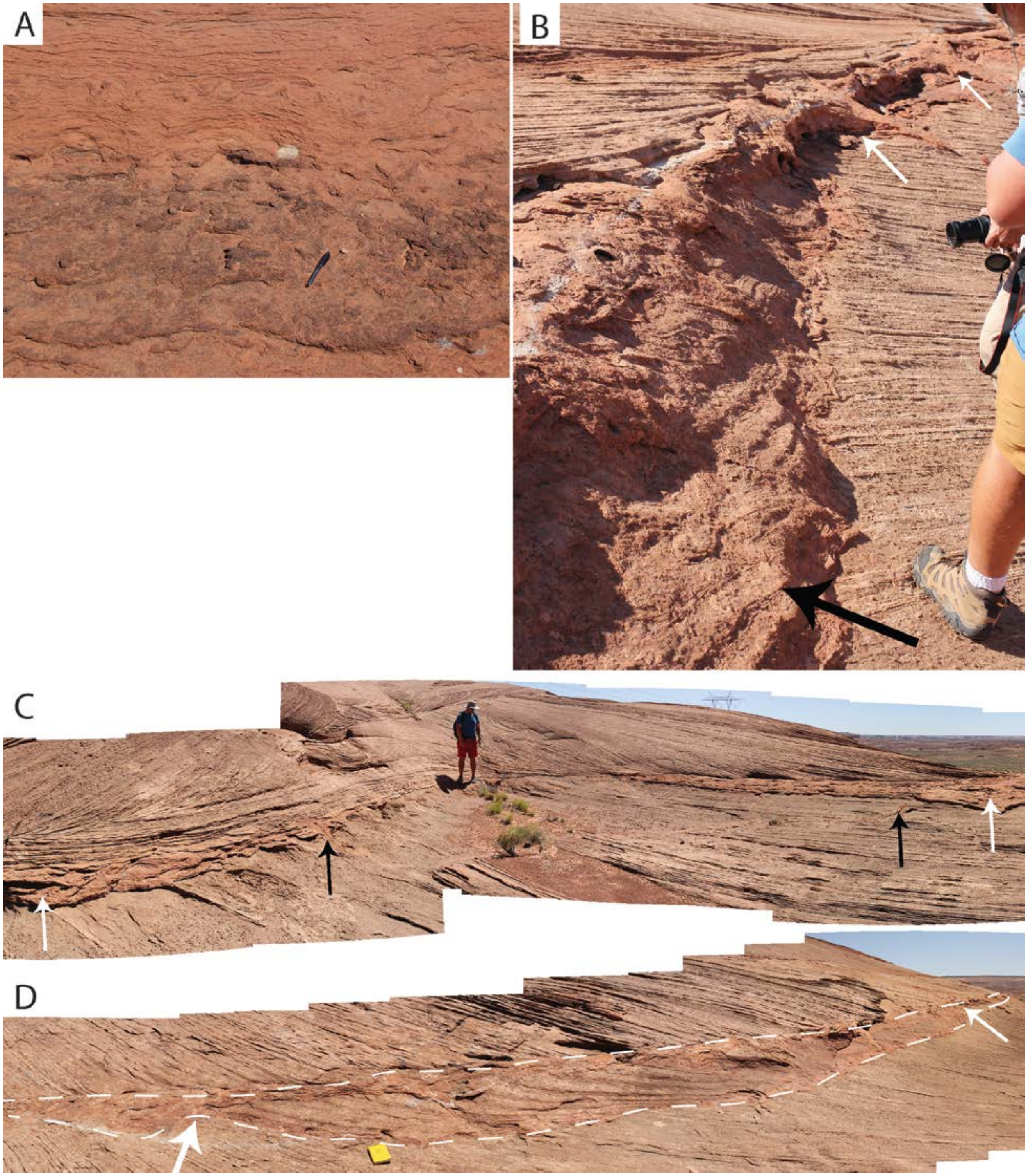

Figure 2 (18.2 cm x 20.8 cm) – A: Contorted sandstone beds interpreted as sabkha deposits. Pen for scale. B: Black arrow points to a thick sabkha deposit along an outcrop-scale bounding surface separating sets of cross-strata. Associated polygonal fractures are visible in the background (white arrows). C: An outcrop-scale bounding surface with variably-thick sabkha deposits. White arrows show the locations with the thickest sabkha deposits. Thickness thins to zero at the person. Black arrows point to polygonal fractures associated with the outcrop-scale bounding surface. From the Ferry Swale outcrop. D: An outcrop-scale bounding surface with a laterally-discontinuous set of cross-strata incorporated into a sabkha deposit (outlined in dashed white lines). White arrows point to the ends of the set of cross-strata. This is an unusual outcrop, and only observed at this location at the Ferry Swale outcrop. Yellow field book for scale.

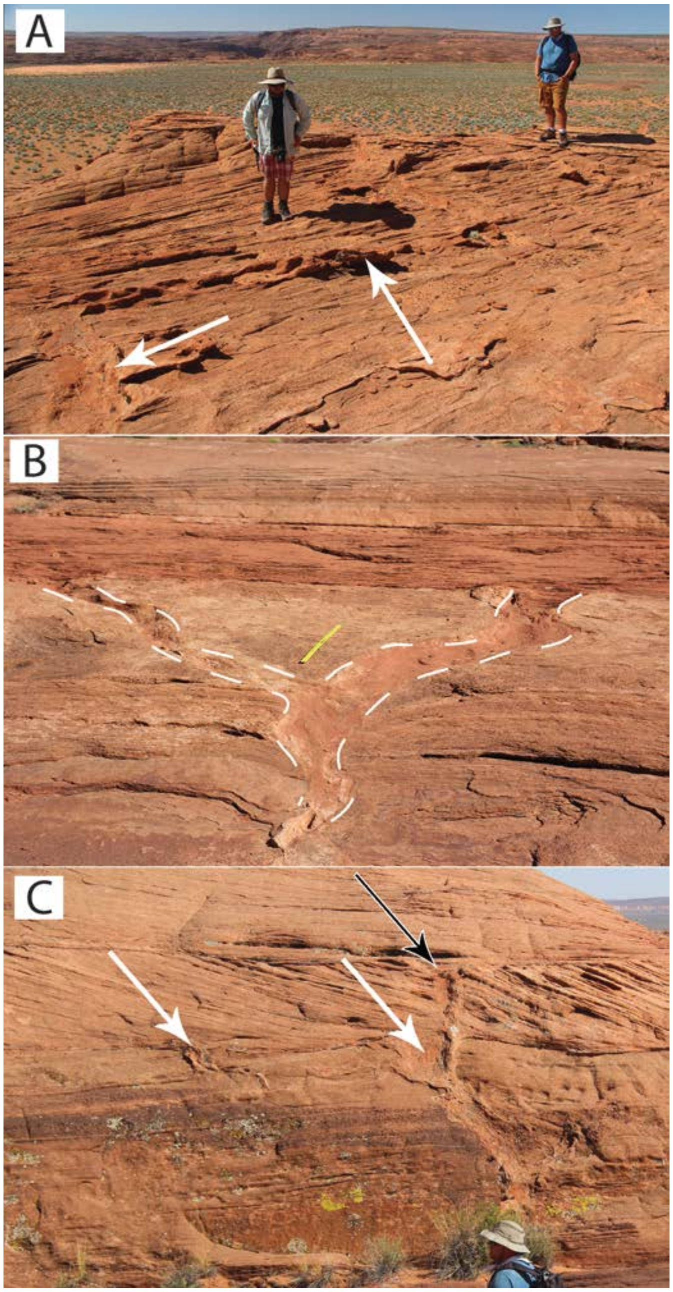

Figure 3 (8.8 cm x 17.0 cm) – Variability of sandstone wedges. A: Topographically-inverted polygonal fractures (white arrows) displaying as raised polygons along a planform exposure of an outcrop-scale bounding surface. B: Two negative relief sandstone wedges intersecting on a shallowly-dipping surface. Wedges are outlined with dashed white lines. Meter stick for scale. C: Sandstone wedges associated with two separate outcrop-scale bounding surfaces at the Ferry

Swale outcrop (the two laterally amalgamating surfaces pointed out in Fig. 9). The black arrow marks the origin of a sandstone wedge at the higher outcrop-scale bounding surface, which cuts beyond the lower outcrop-scale bounding surface. The white arrows mark the origin of sandstone wedges at the lower outcrop-scale bounding surface.

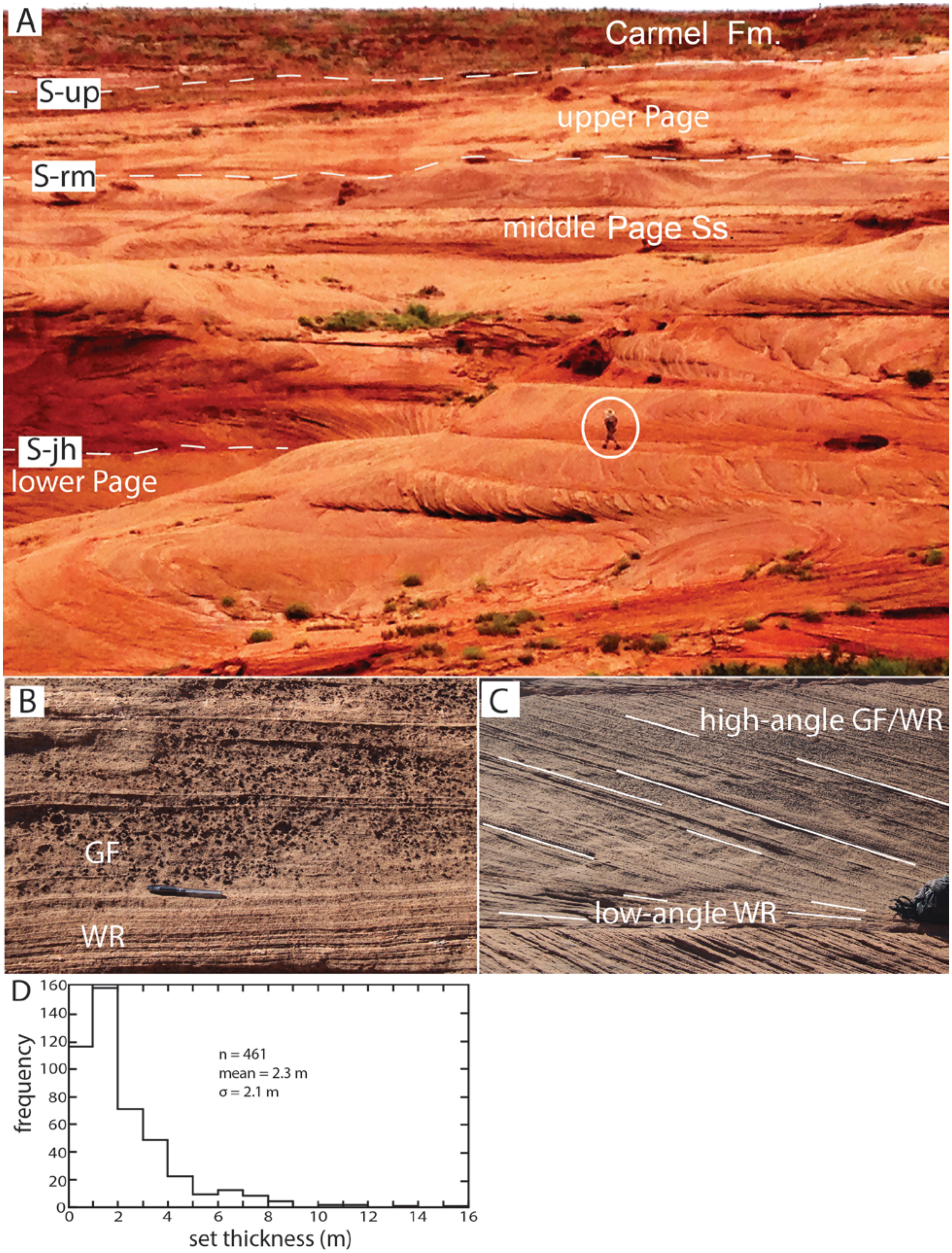
A
Carmel Fm.
S-up
upper Page
S-rm
middle Page Ss
S-jh
lower Page
B
GF
WR
C
high-angle GF/WR
low-angle WR
D
frequency
set thickness (m)
n = 461
mean = 2.3 m
σ = 2.1 m
0
20
40
60
80
100
120
140
160
2
4
6
8
10
12
14
16

Figure 4 (18.2 cm x 23.2 cm) – A: Typical cross-strata composing the lower and middle Page Sandstone. The Carmel Formation is marked on this photograph. Person for scale is circled. Photograph from the Ferry Swale outcrop. Contrast is stretched to highlight set boundaries and cross-strata in the foreground. B: Planform view of typical arrangement of stratification types in sets of the lower and middle Page. Amalgamated grainflow strata (GF) alternate with wind-ripple laminations (WR). Individual grainflow deposits in these packages are generally not identifiable. Rough texture is from concretions frequently associated with the grainflow strata, but not the wind-ripple laminations. Pen for scale. C: High-angle cross-strata composed of alternating textured grainflow strata and wind-ripple laminations transitioning into low-angle wind-ripple strata near the lower bounding surface. Selected cross-strata are mapped in white to highlight this transition. Backpack for scale. D: Histogram showing the high variability of set thicknesses in the lower and middle Page, with respect to the mean.

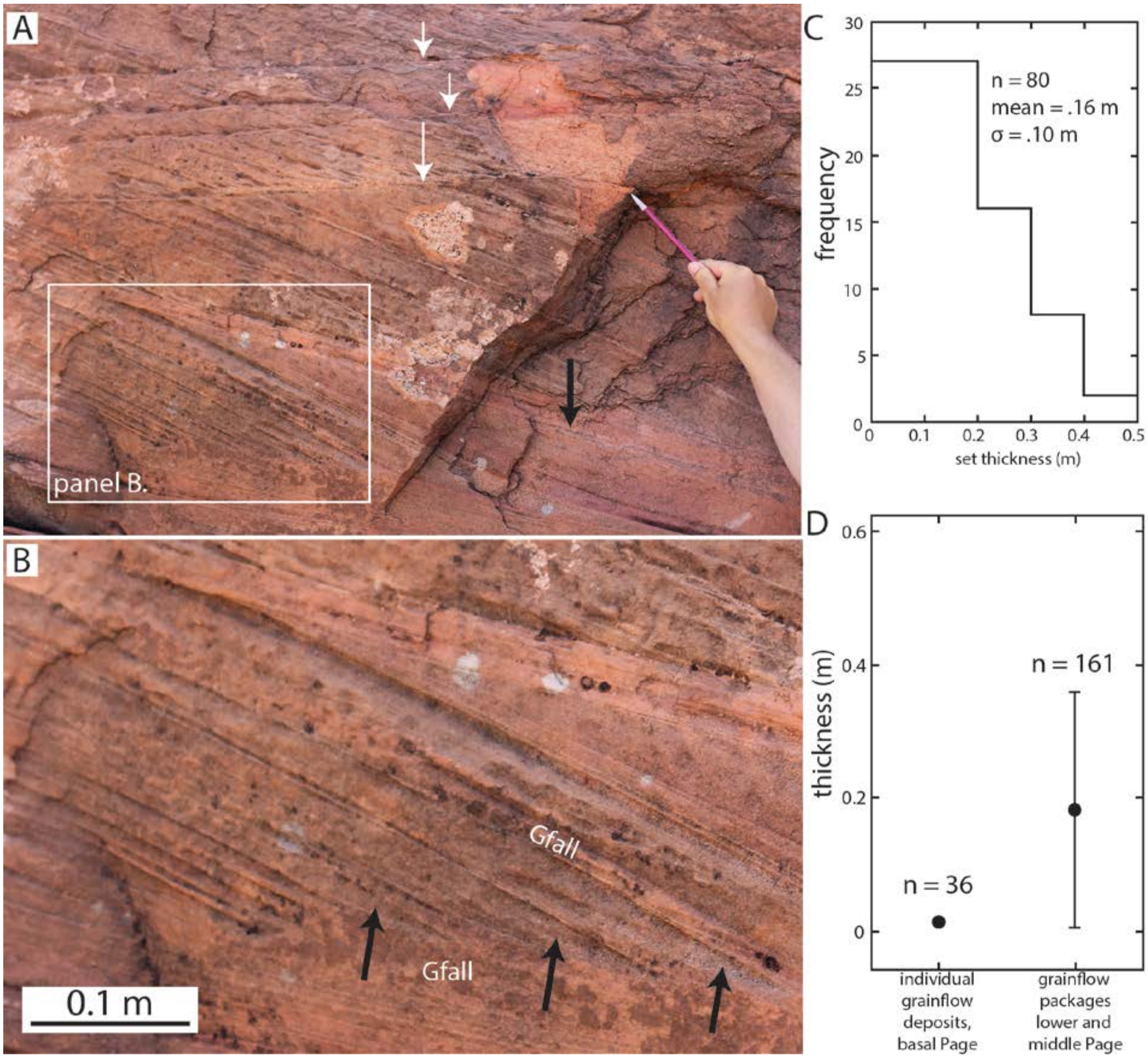

Figure 5 (18.2 cm x 16.3 cm) – A: Close-up of the basal Page sets of cross-strata at the Golf Course outcrop. White arrows point to set bounding surfaces. The black arrow points to a reactivation surface within the set, representing a period of lee face erosion bound by packages representing lee face deposition. The deposit is composed of sets averaging 0.16 m in thickness with individual grainflow deposits separated by thin, tabular beds interpreted as grainfall deposits .The location of panel B is shown with the white box. Pencil for scale. B: A zoom in to part of panel A. Unlike the lower and middle Page sets, individual grainflow deposits are identifiable. Black arrows point to the bottoms of grainflow deposits featuring the characteristic blade shape. Finer-grained grainfall deposits are labeled between and at the base of grainflows, and help distinguish the coarser-grained grainflow deposits. C: Histogram showing the thickness distribution of the thin sets in the basal Page, which is tighter than the distribution of lower and middle Page sets. D: Plot comparing the means (circles) and the standard deviations (bars) of thicknesses of individual grainflow deposits in the basal Page to the lower and middle Page. The standard deviation of the basal Page is within the size of the circle.

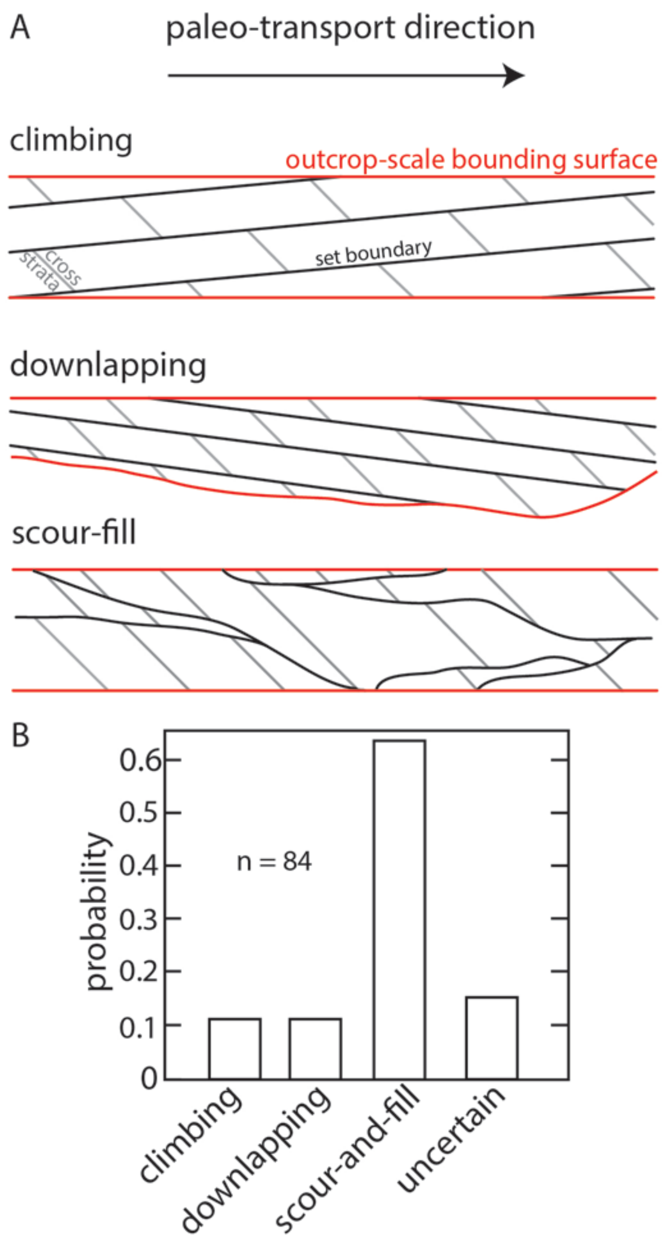

Figure 6 (8.8 cm x 16.2 cm) – A: Set architectures of aeolian cross-strata. The black arrow at the top indicates paleo-transport is from left to right in each case. Cross-strata, set boundaries, and outcrop-scale bounding surfaces are all indicated and consistent between diagrams. Some vertical exaggeration is used to accentuate boundary dips. B: Histogram showing the occurrence of architecture types in the Page Sandstone. The most frequently observed geometry is scour-and-fill. This analysis does not include the western wall of Ferry Swale (Fig. 8) because it is perpendicular to transport direction and does not fit the models in panel A.

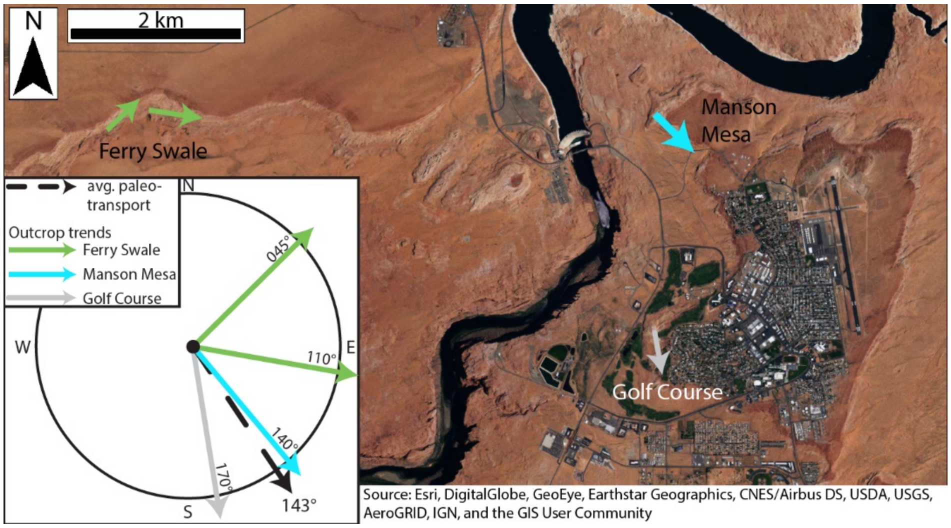

Figure 7 (12.5 cm x 6.9 cm) – Study area near Page, Arizona, marked by the black star in Figure 1. The three outcrops are labeled as Ferry Swale, Manson Mesa, and Golf Course. The associated arrows (green, teal, and white) give the general trends of the outcrop walls, as well as the trends used to create the cross-sections (Figs. 8-9 and S1-S3). In the bottom left corner, the trends of those walls are compared to the average paleo-transport direction within the Page (black dashed arrow; n = 90). The western wall of the Ferry Swale outcrop is perpendicular to the average paleo-transport direction, the eastern wall of the Ferry Swale outcrop is oblique, and the Manson Mesa and Golf Course outcrops are near parallel to paleo-transport.

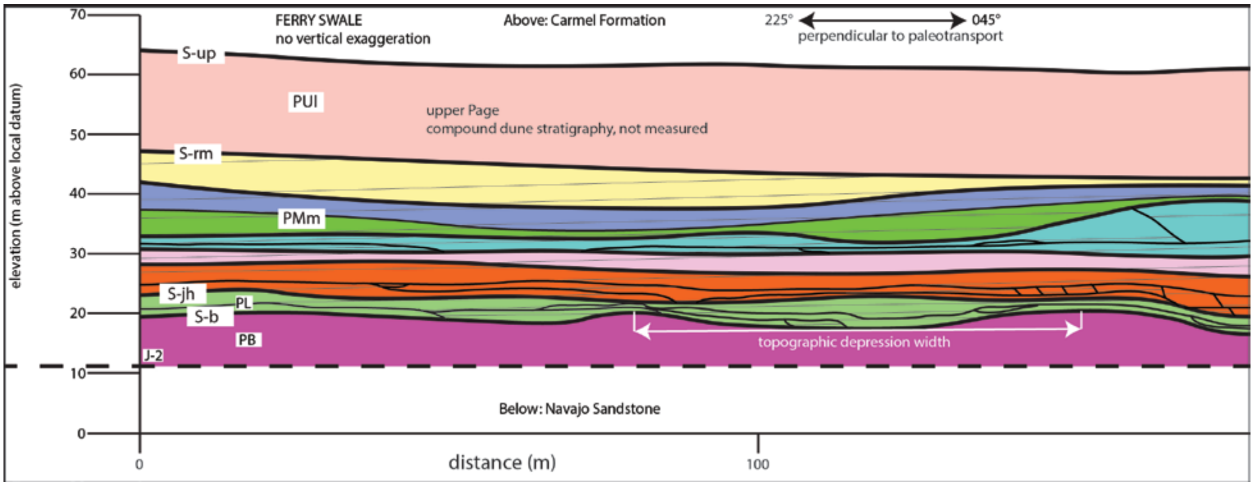

Figure 8 (Landscape, 23 cm x 8.8 cm) – Excerpt of the perpendicular-to-transport wall of the Ferry Swale cross-section (Fig. S1). Location shown in Figure 7. There is no vertical exaggeration. Outcrop-scale bounding surfaces are represented by bold black lines separating outcrop-scale packages of cross-strata of different colors. Black lines within outcrop-scale packages represent cross-set bounding surfaces. Dipping gray lines represent the apparent dip direction of the mean Page transport direction. Colors are used to identify packages of cross-strata separated by outcrop-scale surfaces, and the colors do not necessarily represent correlations between outcrops. The major surfaces are S-up, S-rm, S-jh, and S-b from Havholm et al. (1993). Informal Page units are composed of one to several outcrop-scale packages separated by the outcrop-scale surfaces, and are labeled as in Havholm et al. (1993): PB (basal Page), PL (lower Page), PMm (middle middle Page), and PUl (lower upper Page). The PUl, colored peach, is composed of compound dune deposits, and not studied here. The basal Page (PB) is preserved above the J-2 at this location, and pinches out toward the east (Fig. 9). It is not uncommon for an outcrop-scale package to be composed of one to a few sets, and to vary in number of sets and set thickness laterally. Outcrop-scale bounding surfaces have meters of relief

at this location. The width of a scallop in outcrop-scale bounding surface relief is demonstrated with white double-headed arrow.

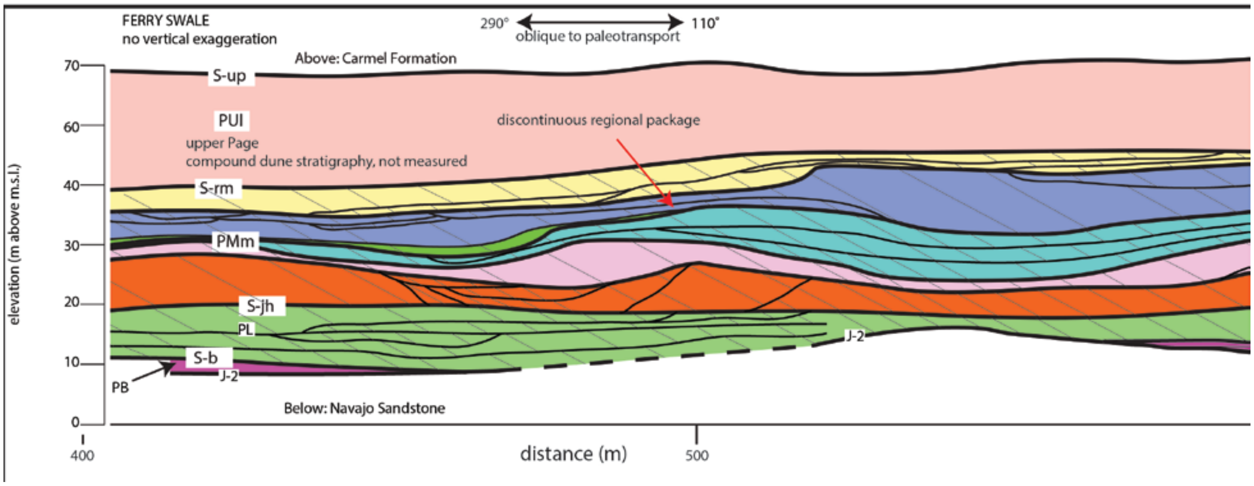

Figure 9 (Landscape, 23 cm x 8.8 cm) – Excerpt of the near oblique-to-transport wall of the Ferry Swale cross-section (Fig. S1). Location shown in Figure 7. Symbology is the same as in Figure 8. Toward the western portion of this section, the basal Page (PB) pinches out along the rising J-2 surface. East of that point, the lower Page sits directly upon the J-2 surface. The location of amalgamation between two outcrop-scale packages is shown with a red arrow. Note the scalloped shape of the outcrop-scale surfaces and the variability in set thickness and discontinuity of sets.

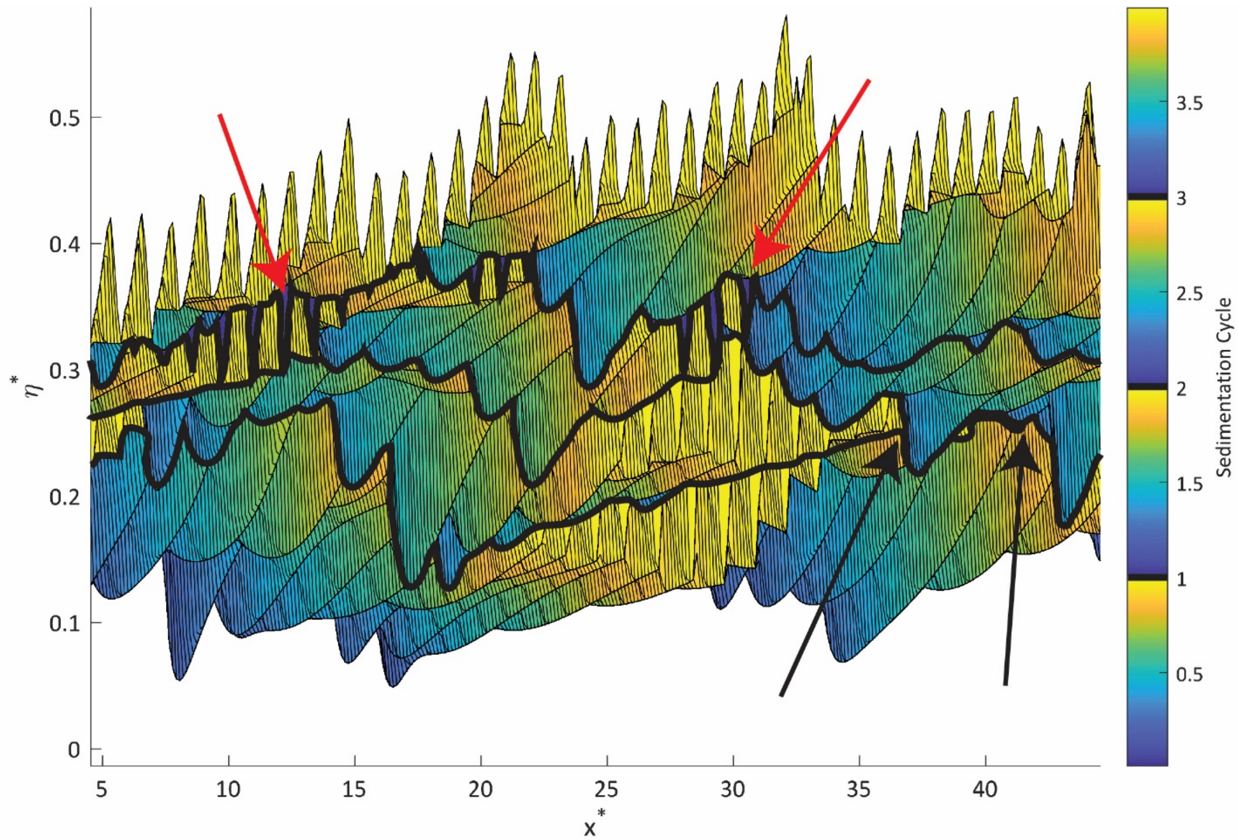

Figure 10 (18.2 cm x 12.2 cm) – Synthetic stratigraphy of Page-type scour-and-fill aeolian accumulation from the model of this paper's companion (Swanson et al. this issue) (model parameters listed in Table S1). Distance and height are expressed in dimensionless values scaled by equilibrium dune wavelength (x-axis) and equilibrium dune height (y-axis). Three water-table highstands are modeled by creating flat, erosional surfaces at progressively higher elevations between episodes of aeolian sedimentation. Following each water-table highstand, a new iteration of the dune field develops. The boldest black domain-scale lines identify three outcrop-scale bounding surfaces associated with the highstands. The next thickest black lines represent set bounding surfaces. Thin black lines represent the cross-strata. Colors within each outcrop-scale package represent the relative accumulation time of the deposit, with dark blue representing earliest-phase accumulation and light yellow representing latest-phase accumulation. Scours tend to preserve strata associated with the cutting dunes and cannibalize early-phase dune field accumulations, hence the absence of significant dark blue accumulations. Red arrows point to examples of a topographically-low dark blue regions representing the earliest-phase fill of antecedent topography. Conceptually, these locations function as an analog to the fill of antecedent J-2 topography by sets accumulated during early phases of the dune field (Fig. 11A-D). In the modeled scenarios, the scour of these accumulations is most frequently performed during later phases of the same sedimentation cycle. Black arrows point to the local amalgamation of two domain-scale bounding surfaces.

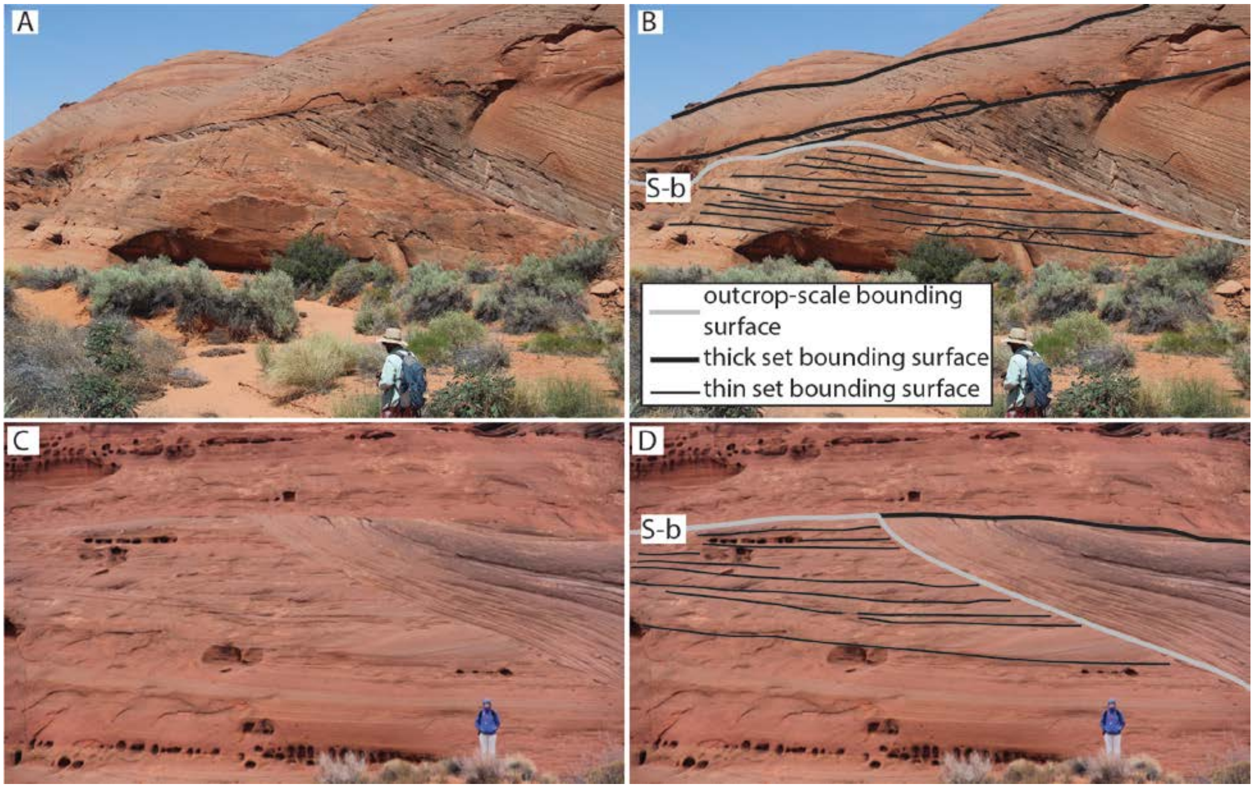

Figure 11 (selected as the color figure for print, 18.2 cm x 11.4 cm) – A: This location is ~100 m north of the Golf Course outcrop. About 5 m of vertical section composed of stacked, relatively thin sets interpreted as basal Page (Fig. 5A-B) truncated by the scour associated with the lower Page. B: Interpreted photograph with the S-b outcrop-scale bounding surface, bounding surfaces of large sets of cross-strata, and bounding surfaces of thin sets of cross-strata mapped. Basal Page is an interpretation, as there is no local exposure of J-2 relief. C: Ferry Swale outcrop. Stacked sets of thin basal Page cross-strata truncated by a scour into the S-b during a later episode of sedimentation. This location is above a 10 m depression in J-2 relief (location in Fig. S1). D: Selected bounding surfaces superimposed onto the photograph. Key is the same as in panel B.

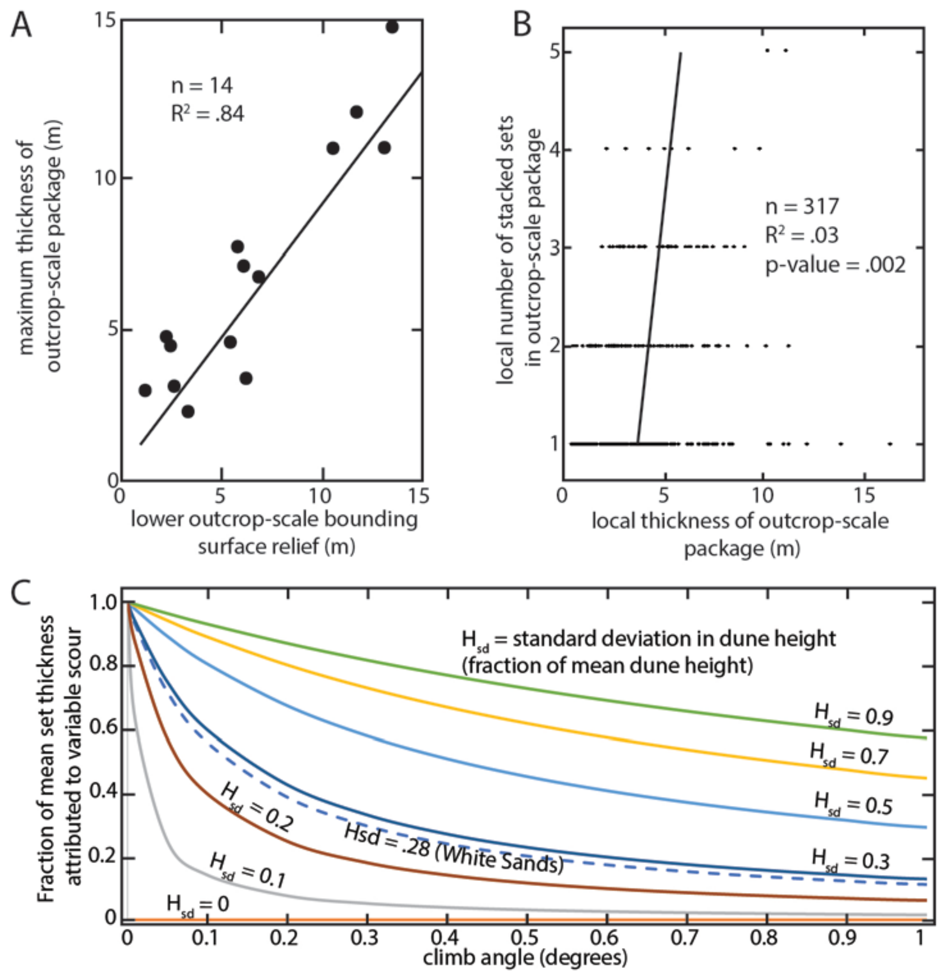

Figure 12 (12.5 cm x 12.8 cm) – A quantification of set architecture composing outcrop-scale packages. Interpretations of the trends are in the Discussion section. A: Total relief along an outcrop-scale bounding surface plotted against the maximum thickness of the outcrop-scale package above the measured surface. The linear fit is significant, indicating that basal relief controls the preserved thickness of the package. B: The number of sets bound by adjacent outcrop-scale bounding surfaces at a given location plotted against the thickness of the package at that same location. The best fit linear trend is not significant, indicating that the number of sets does not control the package thickness. C: Climb angle plotted against the contribution of variable scour to the mean set thickness based on the standard deviation of dune height. Colored lines represent different standard deviations in dune size. At low climb angles, variation in dune height and the subsequent variation in dune scour depths are a dominant control on mean set thickness. An increase in climb angle and/or a decrease in the standard deviation of dune sizes decreases the effect of variable scour on mean set thickness. The fraction of set thicknesses attributed to variable scour depth is calculated for White Sands based on information in (Baitis et al. 2014). Calculated from Bridge and Best (1997) using mean height and celerity values of dunes at White Sands (Baitis et al. 2014).

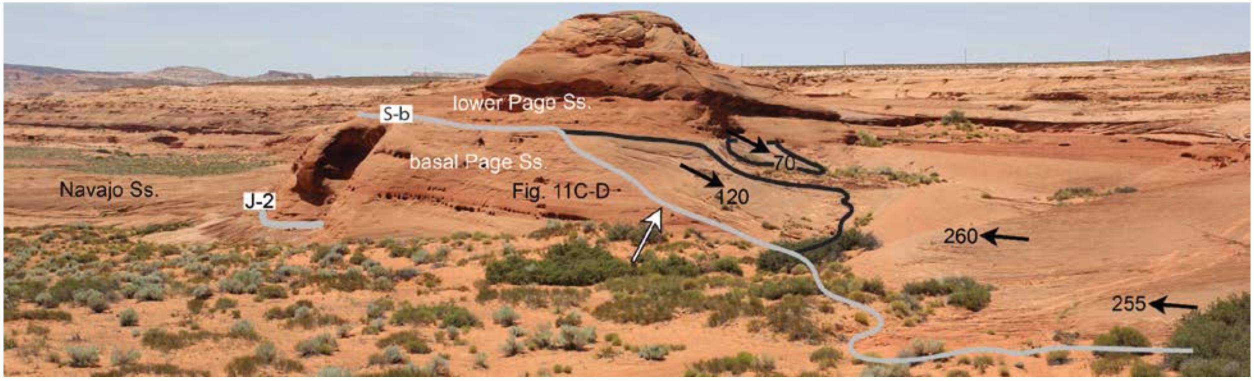

Figure 13 (18.2 cm x 5.5 cm) - Location featuring a topographic depression in the J-2 surface at Ferry Swale, and a subsequent partial scouring of early-phase basal Page accumulations (Fig. S1). The J-2 is mapped with a thick black line, separating the Navajo from the basal Page. Scour surfaces bounding thick grainflow deposits are mapped with thinner black lines. Black arrows show cross-strata dip directions. The stacked thin sets shown in Figure 11C-D are labeled, and are truncated by a scour into the S-b surface dipping towards 120°. Adjacent grainflow strata dip toward 70°. Wind-ripple strata dip towards 260° and 255°, nearly 180° different than the grainflow strata and over 100° different from the regional average transport direction. The center location is interpreted as the deepest portion of a lower Page scour into the S-b surface that partially cannibalized early-phase basal Page.

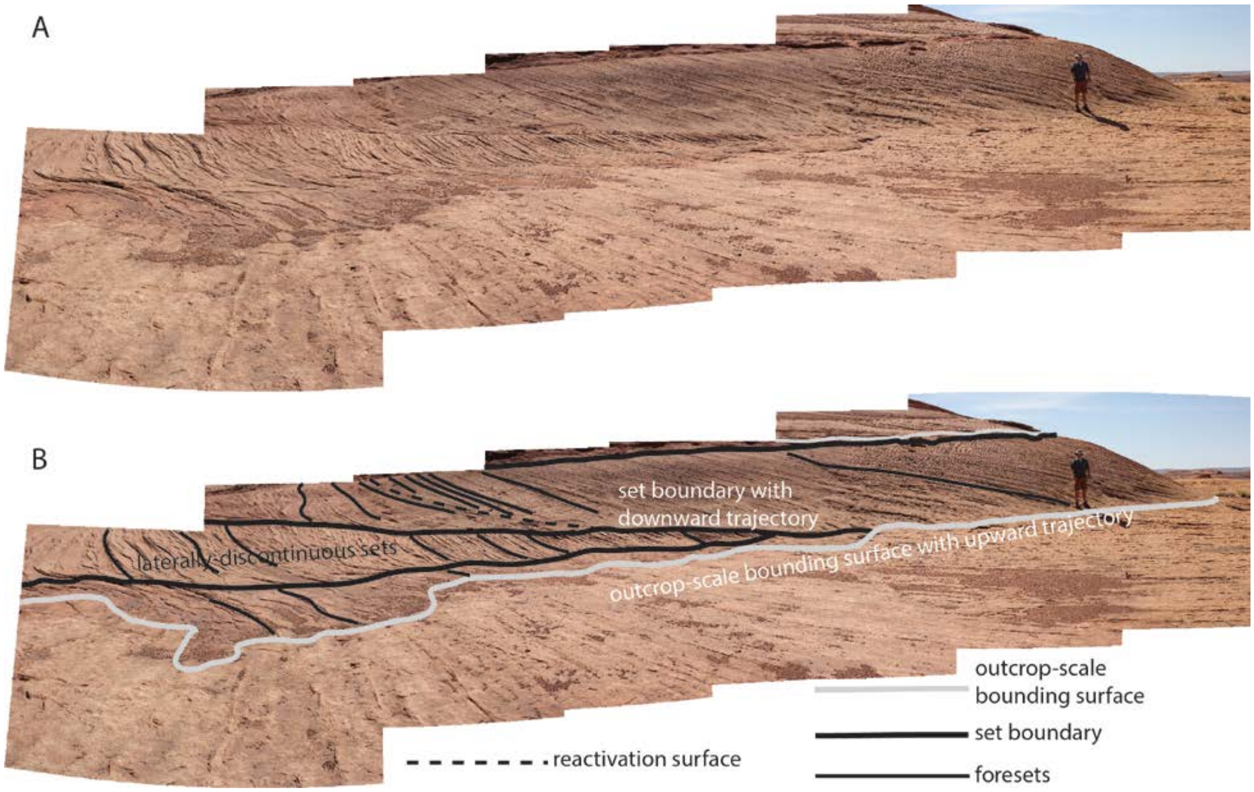

Figure 14 (Landscape, 23.2 cm x 14.6 cm) – A: A set of cross-strata with a lower bounding surface with a downward trajectory, truncating two lower sets of cross-strata within the same outcrop-scale package. B: Interpreted panorama with mapped outcrop-scale bounding surfaces, set bounding surfaces, and a reactivation surface, which represent a period of lee face erosion between periods of lee face deposition. This serves as an example of a scour-and-fill geometry, variable set thickness, and an entire outcrop-scale package defined by only 2-4 stacked sets.

SUPPLEMENTARY MATERIALS FOR:

PRESERVATION OF AUTOGENIC PROCESSES AND ALLOGENIC FORCINGS WITHIN SET-SCALE AEOLIAN ARCHITECTURE II: THE SCOUR-AND-FILL DOMINATED JURASSIC PAGE SANDSTONE, ARIZONA, USA

BENJAMIN T. CARDENAS[1], GARY KOCUREK[1], DAVID MOHRIG[1], TRAVIS SWANSON[2], CORY M. HUGHES[1*], and SARAH C. BROTHERS[1**]

[1]Jackson School of Geosciences, University of Texas at Austin, Austin, TX, USA

[2]Department of Earth Science, Rice University, Houston, TX, USA

[*]now at Department of Geology, Western Washington University, Bellingham, WA, USA

[**]now at National Academies Space Studies Board, Washington, D.C., USA

SUPPLEMENTAL FIGURE CAPTIONS

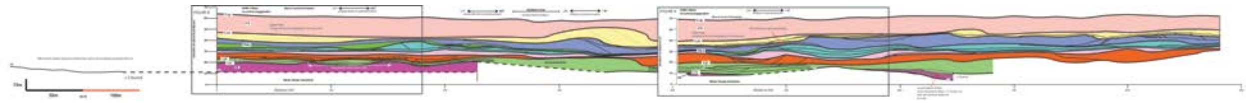

Figure S1 (124.5 cm x 9.5 cm – This will need to be a separate PDF) – Complete cross-section of the Ferry Swale outcrop, including a change in outcrop orientation (location shown in Fig. 7). There is no vertical exaggeration. Outcrop-scale bounding surfaces are represented by bold black lines separating outcrop-scale packages of cross-strata of different colors. The major surfaces among those are labeled S-up, S-rm, S-jh, and S-b (Havholm et al., 1993). Thinner black lines within outcrop-scale packages represent cross-set boundaries. Dipping gray lines represent the apparent dip direction of the mean Page paleotransport direction. The colors used to separate outcrop-scale packages do not represent correlations between outcrop-scale packages shown in Figs. S2 or S3. Informal Page units are composed of one to several outcrop-scale packages separated by the outcrop-scale surfaces, and are labeled as in Havholm et al. (1993): PB (Basal Page), PL (lower Page), PMm (middle middle Page), and PUl (lower upper Page). The PUl, colored peach, is composed of compound dune deposits, and not studied here. Part of the basal Page exposed to the west of the panel and not analyzed here is composed of alternating sabkha and cross-sets. This package of PB is locally preserved above the J-2, and pinches out towards the east. A section of PB exposed towards the east contains the thin, stacked sets shown in Figures 11C-D and 13. It is not uncommon for an outcrop-scale package to be composed of one to a few sets; the pink outcrop-scale package is used as an example. Outcrop-scale bounding surfaces have meters of relief at this location. The width of a scallop in outcrop-scale bounding surface relief is demonstrated with white text and a white double-headed arrow. The outcrop-scale relief of the J-2 surface is shown. The J-2 is at its highest elevation at the westernmost part of the panel, and at its lowest exposed elevation beneath the thin, stacked sets. The locations of Figures 8 and 9 from the main text are shown.

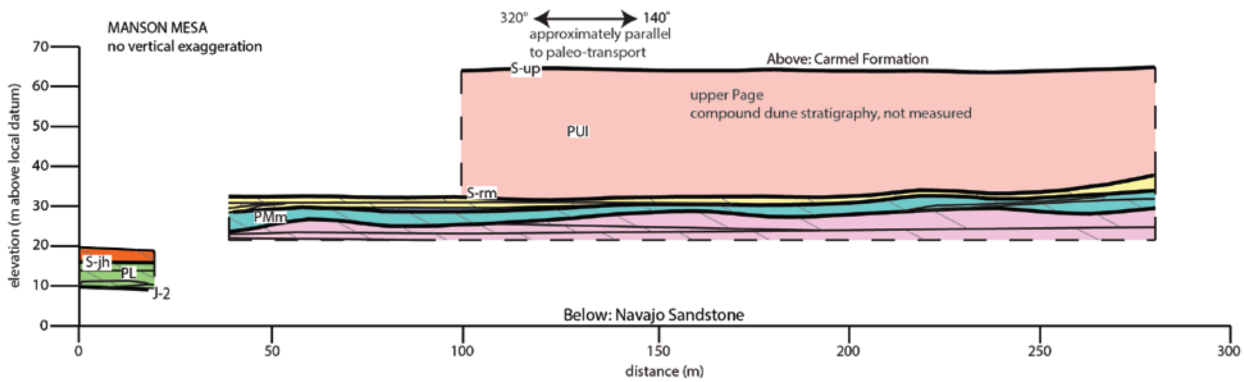

Figure S2 (landscape 23.2 cm x 7.1 cm) – Projected cross-section of the parallel-to-transport Manson Mesa outcrop (location shown in Figure 7). Symbology is the same as in Figure S1. Sets are more laterally continuous at this location, and there is less relief along outcrop-scale surfaces than at Ferry Swale. Exposures of lower Page and S-jh at the left of the panel become vertical towards the right, and eventually are buried, and as such are not continuously surveyed.

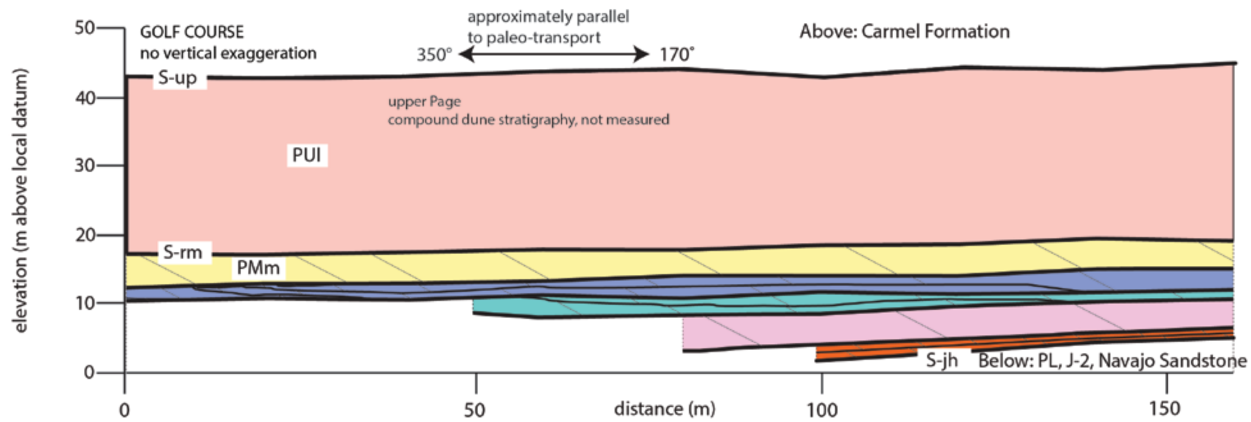

Figure S3 (landscape 23.2 cm x 7.7 cm) - Projected vertical section of the parallel-to-transport Golf Course outcrop (location shown in Figure 7). Symbology is the same as in Figure S1. Like the Manson Mesa outcrop, sets are more continuous and there is less relief along outcrop-scale surfaces compared to Ferry Swale. The vertical wall exposing a package of thin sets is 150 meters north of this location (Fig. 11A-B).

Table S1- Parameters used in the numerical model (Fig. 10). Parameters are discussed in detail in a companion paper (Swanson et al. this issue). Units are arbitrary.

Parameter	Value	Description
A	0.1	Flow blocking parameter
B	3	Flow shoaling parameter
m	1	Meyer-Peter and Müller coefficient
n	1.5	Meyer-Peter and Müller coefficient
E	20	Avalanching coefficient
D	0.2	Diffusivity
p	0.4	Bed porosity
τ_a	0.3	Ambient shear stress
r_a	$1E-4$	Bed aggradation
Δt	1	Model timestep
Δx	10	Node spacing
Δwt	.5	Vertical distance ascended by water table following each deposode