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37

38 **Calderas are subsided volcanic terrains formed by the destructive power of some of the largest**
39 **volcanic eruptions on Earth. Many such depressions globally are today submerged by crater lakes**
40 **or seawater, rendering them less accessible to scientific scrutiny, and hence more complicated to**
41 **monitor during unrest. One of such systems is the restless, partly submerged Campi Flegrei**
42 **caldera (CFc) near Naples in Italy, where escalating hydrothermal activity during ongoing unrest**
43 **demands for an improved understanding of the caldera offshore. Here, we present a high-**
44 **resolution, system-wide mapping of CO₂ degassing encompassing the caldera offshore and**
45 **onshore sectors. Combining vertical seawater profiles with groundwater composition results, we**
46 **obtain a refined delimitation of the anomalous CO₂ degassing zones, finding an overall structural**
47 **control on volatile transport and surface/seafloor discharge. Intense CO₂ degassing on the**
48 **seafloor also causes ocean acidification, sustains diverse chemolithotrophic communities near the**
49 **hydrothermal emissions, and influences the near-vent environment, as demonstrated by**
50 **taxonomic and functional diversity data. Submerged caldera sectors are extremely dynamic**
51 **environments whose volcanological significance and biogeochemistry may have been**
52 **overlooked.**

53 Calderas are volcanic depressions formed by volcano subsidence in response to rapid drainage of
54 magma during large-scale volcanic eruptions¹. Owing to their depressed morphologies, subaerial
55 calderas are recurrently filled with meteoric water forming crater lakes², or invaded by seawater in the
56 offshore of coastal volcanoes. Because submerged calderas escape direct observation from surface, their
57 monitoring poses significant technical challenges to volcanologists, and render these geological objects
58 systematically less well understood than their subaerial equivalents. For example, hydrothermal fluid
59 venting potentially occurs at many, if not all, submarine caldera floors, but information on chemistry
60 and flux of such manifestations remains limited to a relatively small number of better studied systems³.
61 In a few cases, the fluid geochemistry of flooded or partially flooded caldera depressions has been

62 investigated from surveys on the effects of CO₂-rich seepages on shallow marine areas, with notable
63 examples from the Maug Caldera (Northern Mariana Islands)⁴, the harbour of Matupi within the Rabaul
64 caldera (Papua New Guinea)⁵, Levante Bay on the island of Vulcano (Italy) and Mikama Bay in Shikine
65 Island (Japan)^{6,7}. Nevertheless, the extent to which these submersed or semi-submersed caldera systems
66 contribute to the overall (volcano-scale) budget of magma-sourced volatiles remains undetermined.

67 Degassing of CO₂ and reduced volatiles in submarine calderas also supports exceptionally diverse
68 microbial communities, sustained by both photosynthesis and chemosynthesis depending on depth and
69 light penetration⁸. Microbial communities present in the subsurface directly use CO₂ and reduced
70 volatiles (H₂, H₂S, CO, CH₄, Fe²⁺, etc..) of hydrothermal and magmatic origin to sustain
71 chemolithoautotrophy^{8,9}. Similarly high biomass can be found in the venting fluids and in complex
72 biofilm forming at venting sites on the seafloor^{8,10-12}. The organic carbon produced by the microbial
73 communities using volcanic reduced compounds can be exported, together with nutrients, trace metals
74 and reduced compounds, to neighbouring ecosystems^{13,14} and higher trophic levels in the area, thus
75 potentially supporting hot spots of biological activity¹⁵. In addition to this, subsurface microbial
76 communities interacting with volatiles of volcanic and mantle origin have been shown to be able to alter
77 their composition and the quantity of volatiles reaching the atmosphere, potentially impacting volatile
78 budgets in volcanic regions¹⁶⁻¹⁸. In spite of this, the microbial diversity associated with hydrothermal
79 activity in submerged calderas and its biogeochemical contributions to volatile cycling has received
80 limited attention, and the extent of the influence of volatile degassing from submarine calderas to
81 surrounding marine ecosystems remains largely unknown¹⁵.

82 The Campi Flegrei caldera (CFc), on the Neapolitan coast of southern Italy (Fig. 1a), is one of the most
83 densely inhabited and hence hazardous in the world. The caldera has been increasingly monitored in
84 response to a ground deformation, seismic and degassing unrest that, started in 1950-52¹⁹, has continued
85 with alternating vigor until present²⁰, with a remarkable intensification since 2018²¹. Gradual caldera
86 inflation during 4 main unrest episodes (1950-52, 1970-72, 1982-84, and 2005-present) has caused a
87 total net uplift²² of > 4m in the area of maxim uplift (Pozzuoli town), at the centre of the resurgent
88 caldera block²³. Uplift has also been accompanied by intense shallow (< 4 km) volcano-tectonic
89 seismicity²⁴⁻²⁷, especially during two intense uplift phases (known as bradyseismic crises) of 1982-84
90 and 2005-present, and by a visible escalation in hydrothermal degassing activity²⁸.

91 Hydrothermal manifestations within the northern, emerged half of the caldera (Fig. 1) have intensively
92 been studied and monitored, and their temporally changing compositions and fluxes have provided
93 evidence for a sizeable escalation in magma-sourced CO₂ release from the fumaroles^{29,30} and diffusively
94 through degassing soils³¹ and aquifers³². However, the extent to which this CO₂ escalation also extends
95 to the caldera offshore remains comparatively less known³³. This southern, submerged CF caldera sector
96 (Fig. 1) has increasingly been studied using geophysical prospection tool, allowing for a refined

97 understanding of caldera floor morphology³⁴, and of its geological³⁵ and structural^{35–39} setting and
98 evolution, including the effects of relative sea water changes and their interplay with human settling
99 and infrastructures dating back the roman time^{40,41}. Marine geology prospecting has also provided
100 observational evidence for an intense degassing of magmatic fluids on the submerged caldera floor³⁴,
101 but direct measurements on the chemistry of such submarine hydrothermal vents have been limited to
102 a few cases³³, and the impact of vent emissions on the chemistry and biology of the overlying seawater
103 column and on the surrounding ecosystems is poorly understood. Studies on the microbiology and
104 ecosystems impact of the CFC degassing are limited to a preliminary description of the microbial
105 diversity associated with the fluids and biofilms of some shallow-water hydrothermal vents in the
106 shallower parts of the caldera¹² and an investigation of the sedimentary meiofauna diversity associated
107 with a single vent⁴².

108 **Results**

109 **Offshore pCO₂**

110 We conducted a systematic survey across the Pozzuoli Bay (Fig. 1b), during which we characterized
111 degassing activity in the CFC offshore by measuring the CO₂ partial pressure (pCO₂) in the seawater
112 column at 273 vertical profiles (see Methods and Supplementary Table 1). Our large dataset (520
113 punctual measurements) constrain the mean and median pCO₂ values at 574 and 505 μatm, respectively
114 (± 1 standard deviation of 502 μatm). This is well above the typical atmospheric equilibrium pCO₂ value
115 of 420 μatm⁴³, indicating the existence of an additional CO₂ source in excess to atmospheric CO₂.

116 We identify a large spatial heterogeneity in pCO₂ values across the bay (Fig. 2), suggesting that such
117 excess CO₂ is not uniformly distributed, but rather focused in specific areas. Approximately 80% of
118 seawater measurements fall within a relatively narrow range, from approximately 376 μatm to 600 μatm
119 (Figs. 2a to 5). However, significantly higher concentrations are observed at a number of sites, revealing
120 a "hotspot" distribution. These localized, high-concentration zones cluster in 5 main areas where CO₂-
121 rich hydrothermal vents have been reported in the seafloor (Fig. 2a, 3):

122 Secca delle Fumose (SdF): This area is known since antiquity for the presence of hydrothermal CO₂
123 vents^{33,44} distributed along pillars of Roman age of unknown use (Supplementary Figures 1a, b). This
124 site exhibits the most extreme pCO₂ conditions observed, and was therefore selected for coupling pCO₂
125 measurements with microbial diversity analysis in a series of water column transects (Fig. 2b). A
126 vertical profile was conducted in the SdF area, right above the SdF vents, each consisting of pCO₂
127 measurements at 2, 6, and 10 m depth (Fig. 3a and Supplementary Table 2). This profile demonstrates
128 a steep CO₂ vertical gradient along the water column which leads to a rapid decline in seawater pH. At
129 2 m, pCO₂ ranges from 500–520 μatm (pH = 8.02), suggesting only moderate enrichment relative to
130 ambient seawater. At 6 m, pCO₂ increases substantially, reaching 1400–1600 μatm (pH = 6.94). The

131 highest values are recorded at 9-10 m, where pCO₂ peaks at 6800 to 9000 μatm (pH = 5.77). In
132 conjunction with nearby profiles, these observations identify a clear CO₂ plume originating from the
133 seafloor, and propagating upward along the seawater column (Fig. 3a, 3b). Orange and green colored
134 biofilm are present in the area of focused venting (Supplementary Figure 2) while white biofilms are
135 present in the area of lower temperature diffuse degassing (Supplementary Figures 3 and 4).

136 Baia Castle beach: Here, pCO₂ levels are elevated but remain substantially lower than those recorded at
137 SdF (Figs. 2-3b). Owing to the shallow seafloor, only a single depth measurement was possible. pCO₂
138 measured at 3 m reached a maximum of 849 μatm. Although this value indicates moderate CO₂
139 enrichment, free gas bubbles were observed rising from the substrate at this location, confirming the
140 presence of localized venting activity. The limited water depth prevented assessment of a vertical
141 gradient, yet the measurement and visual observations together characterize Baia as a site influenced
142 by mild but detectable CO₂ seepage.

143 Capo Miseno: Bubbling activity has systematically been reported in the Capo Miseno area, both at the
144 seafloor and at the air-seawater interface, and was clearly visible during our campaign (Supplementary
145 Figures S1c, d). Here, our results identify two spatially distinct CO₂-emitting hotspots. The first occurs
146 within the Gulf of Miseno, between Isola Pennata and Punta Sarparella (Fig. 3c). In this area, 22 vertical
147 profiles were conducted, with measurements at 1–2 m and 7–10 m depth (water depth reached ~20 m
148 at stations near the gulf entrance). pCO₂ values increase with depth and are uniformly elevated ranging
149 from 540 to 1700 μatm across profiles. A second CO₂-rich spot is identified south of Capo Miseno, in
150 the vicinity of the Faro (Fig. 3c and Supplementary Figure 5). Free gas emissions in this site were
151 visibly more intense than in the gulf, with continuous bubble release reaching the surface, indicating
152 stronger localized venting. Here, a detailed profile was completed with measurements at two depths:
153 pCO₂ at 1–2 m is ~590 μatm, while at 7–8 m it increases substantially, ranging from 610 to 2300 μatm.
154 Although more localized than the first hotspot, this site similarly exhibits a pronounced increase in
155 pCO₂ with depth.

156 Accademia-Bagnoli area: A total of 62 vertical profiles are conducted along the Accademia and Bagnoli
157 coastlines, representing the most spatially extensive survey area in this study. Profiles closest to the
158 shoreline—where the highest pCO₂ values are recorded—consist of a single measurement at 1–2 m
159 depth due to the shallow seafloor. In contrast, offshore profiles include two or three depth layers (1–2
160 m, 3–7 m, and 9–15 m), depending on local bathymetry. The highest pCO₂ values in this area are
161 measured along the Bagnoli coastline, particularly in proximity to the Pontile Nord. Here, pCO₂ reached
162 a maximum of 4420 μatm, with additional peaks of 882 and 2450 μatm recorded at nearby points. Free
163 gas bubbles are frequently observed rising from the seabed, confirming the presence of active degassing.
164 In the Accademia sector, pCO₂ levels are lower (576–732 μatm) but still markedly elevated relative to
165 typical atmospheric equilibrium concentrations (Fig. 2, 3b).

166 Nisida: The Nisida island offshore is a significant CO₂-emitting spot within the Gulf of Pozzuoli (Figs.
167 2-3b). Elevated pCO₂ levels are detected at the base of the peninsula, particularly in the area facing the
168 Lega Navale di Pozzuoli and the Dente di Cane rocky outcrop. A total of 35 vertical profiles are
169 conducted across the area, allowing characterization of both nearshore and offshore variability. At 1–2
170 m depth, pCO₂ ranges from 800 to 1100 μatm, indicating a localized but persistent CO₂ excess reactive
171 to air-saturated seawater.

172 Measurements of seawater pCO₂ were not possible in three sectors of the Pozzuoli Bay due to access
173 limitations and maritime regulations (Figure 1b). These include (i) the interior of the Pozzuoli harbour,
174 where active venting has been reported but could not be surveyed because of restrictions on boat
175 operations; (ii) the high-traffic navigation corridor used by local ferries, which is excluded from
176 scientific access for safety reasons; and (iii) an area of approximately 0.723 m² southwest the SdF
177 occupied by a mytilid aquaculture facility, where safe operation of powerboats is not feasible. Although
178 these sectors were interpolated in the spatial analysis, they may host additional CO₂ sources, implying
179 that our estimates of offshore degassing are likely conservative.

180 **Onshore pCO₂**

181 CO₂ degassing in the onshore CFC sector is evidenced by high pCO₂ values (n = 83) measured in
182 groundwater samples. In this study, pCO₂ was recalculated in μatm using data previously reported in
183 the literature³² (Methods, and Fig. 1b, 4). The extremely high values encountered reflect the large,
184 widespread interaction of the groundwater system with deep-rising magmatic CO₂, sourced by the
185 underlying magma plumbing system⁴⁵. Groundwater pCO₂ displays a broad range, from 317 μatm to
186 867961 μatm. The mean pCO₂ is 116512±208011 μatm (mean ± standard deviation), with a median of
187 21429 μatm. 67% of all land-based measurements exceed 10000 μatm, and 29% exceed 50,000 μatm.
188 CO₂ enrichment is the largest in a wide degassing area centered on Solfatara crater, the largest degassing
189 structure in the caldera onshore²⁸, and stretching from Pozzuoli town to the west to Agnano-Bagnoli to
190 the east (Fig. 4). A second, less prominent CO₂ degassing zone is identified on the west, in the Baia-
191 Averno-Monte Nuovo area (Fig. 4).

192 **Microbial diversity in the hydrothermal vent environment**

193 Geochemical pCO₂ surveys were coupled with fluid, biofilm, and seawater sampling to investigate the
194 effects of hydrothermal venting on microbial diversity. Between June and July 2023, divers collected
195 hydrothermal fluids (73 °C) and biofilm samples directly from the SdF vent at 10 m depth, as well as
196 biofilms associated with lower-temperature (30–40 °C) diffuse venting within the same hydrothermal
197 area. In October 2024, seawater column samples were collected along N–S and E–W transects centered
198 on the SdF vent (n = 9 water column stations, for a total of 21 samples at different depths; Fig. 2b).
199 These samples are part of a broader framework of biological sampling efforts aimed at characterizing
200 microbial diversity in the Gulf of Pozzuoli.

201 Shotgun metagenomic sequencing generated 6,997 million reads across all samples (a mean of 259
202 million reads per sample), of which 99% passed quality filtering. Of these, 802 million paired reads
203 (23%) were taxonomically annotated. The vent fluid sample (SF1D13_F2) was dominated by
204 Aquificota (53.99% of annotated sequences), largely represented by the Hydrogenothermaceae family
205 (50.75%; Fig. 6). Biofilm samples were dominated by a limited number of phyla but showed marked
206 compositional differences. The biofilm collected at the SdF vent orifice (SF1D13_B1) was strongly
207 dominated by Cyanobacteriota (36.55%), whereas a second biofilm from the same vent (SF1D13_B2)
208 was enriched in Aquificota (22.89%), with Hydrogenothermaceae accounting for 21.46% of the
209 community. Microbial white mats collected from lower-temperature diffuse emission sites (SF2A_B1,
210 SF2B_B1, SF2C_B1) shared the same dominant phyla but exhibited a clear compositional shift across
211 samples. Pseudomonadota decreased from 32.23% (SF2A_B1) to 12.08% (SF2C_B1), while
212 Campylobacterota increased from 2.91% to 15.45% of total reads. Within Campylobacterota,
213 Sulfurovaceae ranked among the most abundant families in SF2B_B1 (2.69%) and SF2C_B1 (1.99%),
214 whereas Thiovulaceae peaked in SF2C_B1 (6.33%). These all represent known groups of thermophilic
215 and mesophilic chemolithoautotrophic microorganisms capable of hydrogen, sulfide and thiosulfate
216 oxidation at the expenses of nitrate or oxygen⁴⁶⁻⁴⁸.

217 Seawater samples displayed a broadly homogeneous background community at the phylum level,
218 dominated by Pseudomonadota (mean ~47.45%, range 29.65–72.85%) and Bacteroidota (mean
219 ~13.53%, range 3.36–22.66%; Fig. 6), while Cyanobacteriota and Verrucomicrobiota remained minor
220 components. At the family level, Alteromonadaceae, Pelagibacteraceae, Flavobacteriaceae,
221 Paracoccaceae, and Halieaceae were consistently among the most abundant. Some samples showed
222 localized enrichments in less ubiquitous families (*e.g.*, Halomonadaceae, Vibrionaceae), including clear
223 outliers such as T7D1, where Pseudoalteromonadaceae reached 14.49%. These sequences represent
224 well known coastal water column microbial groups involved in oxygenic photosynthesis (sequences
225 belonging to the Cyanobacterota) or in organic matter cycling^{49,50}.

226 Compositional differences among samples were assessed using beta diversity analysis based on
227 weighted Jaccard dissimilarity (Fig. 7a). Seawater samples clearly separated from biofilms and
228 hydrothermal fluids along the main PCoA axis, which captured 36.5% of the total variance. Within
229 seawater samples, separation reflected the degree of hydrothermal influence, primarily associated with
230 pH variations along the first and second ordination axes, with PCoA1 correlating with pH (Fig. 7b,
231 Pearson correlation, $r^2 = 0.549$, $p = 0.012$). This separation was driven by microbial groups typically
232 associated with hydrothermal emissions, which were abundant in vent fluids and biofilms (Fig. 8).
233 These families constituted up to 50% of annotated sequences in hydrothermal-associated samples (Fig.
234 8a), but remained rare in seawater (<0.1% on average), showing systematic increases with decreasing
235 distance from the vent and increasing depth, consistent with dilution during plume mixing (Fig. 8b).

236 The distribution of genes encoding key biogeochemical functions mirrored the taxonomic patterns (Fig.
237 9). PCoA of functional gene composition based on weighted Jaccard dissimilarity showed a clear
238 separation between seawater and hydrothermal-associated samples along the main axis, accounting for
239 51.3% of the variance (Fig. 9a) when using all the biogeochemical genes, while the second axis
240 explained only 9.3%. Genes of the Calvin–Benson–Bassham (CBB) cycle and the 3-hydroxypropionate
241 (3HB) bicycle were widespread across all samples (Fig. 9b). In contrast, genes encoding the reductive
242 tricarboxylic acid (rTCA) cycle and the Wood–Ljungdahl (WL) pathway were enriched in vent fluids
243 and biofilms, reflecting their association with facultative microaerophilic and anaerobic
244 microorganisms typical of hydrothermal systems. Nevertheless, these genes were also detected in
245 seawater samples, consistent with plume-mediated transport of vent-associated microorganisms and
246 metabolic potential (Fig. 9b).

247 Vent fluids and biofilm samples shared a pronounced sulfur cycling potential, being enriched in genes
248 involved in thiosulfate disproportionation, sulfur, sulfite, and sulfate oxidation, as well as sulfate
249 reduction (Fig. 9c). Sulfur cycling genes were also detected in seawater, albeit at lower abundances,
250 particularly those involved in thiosulfate disproportionation and sulfide oxidation. While present across
251 all sample types, genes associated with thiosulfate oxidation were relatively enriched in seawater
252 compared to SdF biofilms, suggesting that this metabolism is favored under the more oxidizing
253 conditions characteristic of plume mixing zones.

254 Discussion

255 **Caldera offshores as potential sites of intense hydrothermal degassing.** Calderas on Earth
256 occasionally become restless in what is referred to as a caldera unrest^{51,52}. In addition to ground uplift
257 and heightened seismicity, commonly observed signs of caldera unrest include macroscopic changes in
258 hydrothermal degassing (*e.g.*, enlarging of exhaling zones and increasing volatile fluxes) and more
259 subtle compositional changes in surface hydrothermal manifestations, whose detection and
260 interpretation requires careful scrutiny by volcano geochemists^{29,28,53,54}. However, manifestations
261 (fumaroles, soils, mud pools, groundwaters) normally targeted by research/monitoring^{55,56} are typically
262 limited to the emerged (subaerial) portions of calderas, leaving their offshore sectors frequently
263 unexplored. At calderas that are in (large) part submerged (for example, roughly one third of the CFc
264 lies underwater; Fig. 1), ignoring hydrothermal venting on the seafloor can cause a (potentially severe)
265 underestimation of caldera degassing, eventually also limiting our ability to fully capture dynamics and
266 temporal evolution of unrest.

267 Our results here (Figs. 2, 3) indicate that large portions of the CFc offshore are characterized by above-
268 atmospheric pCO₂ values (Fig. 5). In the 5 high pCO₂ zones described above (see results), vigorous
269 bubbling activity is visible, both on the seafloor and at the air-seawater interface. These results suggest

270 that the venting of CO₂-rich fluids is recurrent and widespread in several segments of the submerged
271 CFC offshore. We anticipate that submarine hydrothermal venting may be a recurrent process at many
272 (most) partly to totally submerged calderas worldwide. Caldera-hosted degassing systems span different
273 geodynamic settings, and a wide spectrum of depths and environmental contexts, from shallow tropical
274 coral reefs, subtropical and temperate nearshore habitats – including sandy bottoms, seagrasses, and
275 algal rocky reefs^{4-7,57,58} – to deepsea arc/back-arc calderas like those in the Tonga-Kermadec, Eolian
276 Islands and the Mariana arcs⁵⁹⁻⁶¹. This environmental heterogeneity highlights that submarine calderas
277 can influence diverse marine ecosystems across latitudes and settings, with implications of caldera
278 degassing extending well beyond any single site. The diversity of emission forms and environments
279 makes extrapolation of the few available data very difficult. Hence, the global CO₂ output from this
280 form of volcanic degassing, which is missing from current inventories⁶²⁻⁶⁵, needs to be quantified by the
281 implementation of a more systematic and widespread observational program.

282 **A structural control on CO₂ seafloor discharge.** The seawater observations here, integrated with
283 previous results for the groundwater system inland³², allow a system-wide mapping of CO₂ degassing
284 that includes both the CFC onshore and offshore (Fig. 4). This is the first map of its kind at an active
285 caldera (at least to our knowledge), and offers a more complete view of CO₂ degassing extents and
286 pathways at CFC than possible with terrestrial observations alone.

287 In the central, eastern portion of the map, we identify an overall land-to-sea continuity in high CO₂
288 degassing in a sector (sector 1 in Fig 4a) that extends from Pozzuoli-Agnano-Bagnoli inland to their
289 adjacent offshores. This large high-CO₂ anomaly is centered on, but wider than, the Solfatara crater⁶⁶, a
290 tuff cone/maar-diatreme structure^{67,68} formed circa 3.9 ka BP, and representing the largest diffuse
291 degassing structure in the CFC^{28,31,69}. Our results here (Fig. 4) indicate that the CO₂ anomaly not only
292 extends west (toward Pozzuoli), north (toward Astroni crater), east (Bagnoli) and north-east (Agnano)
293 of Solfatara crater, as previously noted³², but also to the south in the Bagnoli-Accademia offshore (Fig.
294 3-4). It is also worth noting that the seawater CO₂ anomaly, stretching from Accademia to Bagnoli (Fig.
295 3a and 4), roughly corresponds to the area where an incipient fault system⁷⁰, where more than 50% of
296 the post-2022 seismicity is thought to have been concentrated (Fig. 1,4), would intersect the seafloor
297 (Fig. 4b). The Accademia geodetic anomaly⁷¹, an area of reduced (slower) uplift that manifested since
298 2021, causing a local deviation from the typical bell-shaped deformation pattern, is located north-
299 northwest of this area. We hence interpret the Pozzuoli-Accademia-Bagnoli seawater anomaly as due
300 to a combination of (i) direct CO₂ contribution from the incipient fault system and surrounding highly
301 fractured zones, affected by recurrent seismicity, and (ii) discharge into seawater of CO₂-rich aqueous
302 fluids delivered by the aquifer system inland (Fig 4b). This latter process is perhaps also responsible for
303 the seawater pCO₂ anomaly in the Baia Castle beach area (Fig. 2), where groundwaters in the adjacent
304 inland sector are similarly CO₂-rich (sector 2 in Fig 4a).

305 In addition to delimiting the extension into seawater of previously known anomalous CO₂ degassing
306 zones (sectors 1 and 2 above), our results also identify areas (sectors 3-5) of degassing that would
307 otherwise have escaped detection from terrestrial observations alone.

308 Our caldera-wide survey identifies the SdF CO₂ plume well (Fig. 2, 3). Hydrothermal venting of CO₂-
309 rich fluids⁴⁴ at SdF has been known for centuries⁷², and the CO₂/heat output from this manifestation has
310 already been characterized in a dedicated study³³. Our map here (Fig. 4) indicates that the SdF CO₂
311 anomaly is in correspondence with trending faults that have been active^{73,24,27,74} during the ongoing
312 unrest (see Fig. 4b) and before/during the Monte Nuovo eruption, when this fault system fed the eruptive
313 vent^{33,75}, destroying part of the Roman structures³⁴. These N-S faults are part of a wider fault system in
314 the caldera offshore²⁷ that correspond to the bordering faults of the caldera resurgent block³⁹. Vigorous
315 CO₂ venting at SdF confirms that this fault system is still active. It is therefore likely that the southern
316 segment of the resurgent block bordering faults, not investigated in this study (seawater depth extends
317 to more than -80 m), is also affected by active hydrothermal venting.

318 Finally, the high CO₂ anomalies in the Nisida and Capo Miseno areas (sectors 4 and 5 in Fig. 4a) are
319 found at the seawater intersection of the outer caldera rim bordering faults. These faults formed during
320 the 15 ka old Tufo Giallo Napoletano (TFG) caldera-forming eruption⁷⁶. However, at both Nisida and
321 Capo Miseno, these caldera rims are home to eruptive vents that were active at the end of the third
322 epoch of the CF volcanism. Hence, although volcanism at CFc has mainly clustered in the caldera center
323 during the last 5.5 ka⁷⁷, our results suggest that the caldera outer bordering faults may be degassing and
324 active more than groundwater results inland would suggest.

325 **Caldera-scale submarine CO₂ degassing: significant or marginal?** Our results set the basis for a
326 first, very preliminary quantification of the total CO₂ output sustained by the CFc offshore. To this aim,
327 we rely upon the pCO₂ results illustrated in the cross-section of Fig. 3b that, stretching in a ESE-WNW
328 direction across the Pozzuoli bay, is roughly perpendicular to the dominant NE-SW seawater current
329 direction (see below, Supplementary Figure 6).

330 We initially use a software⁷⁸ specifically designed to solve the seawater carbon system to convert the
331 pCO₂ values (of Fig. 3a; range, 423-8029 μatm) into Total Dissolved Inorganic Carbon (TDIC).
332 Calculations are run using average salinity, temperature and alkalinity of respectively 37.9‰, 19.7 °C
333 and 2600 μmol/kg, and output TDIC values of 2280-2830 μmol/kg.

334 Next, to quantify the TDIC contribution related to hydrothermal seepage, the background (ambient)
335 seawater TDIC must be subtracted. This includes any atmospheric and biogenic carbon contribution to
336 the seawater system. Our pCO₂ population is manifestly polymodal (Fig. 5), and is best fit (AIC, Akaike
337 Information Criterion: $6.47 \cdot 10^3$; Bayesian Information Criterion, BIC: $6.51 \cdot 10^3$) by a mixture of 3 sub-

338 populations each with gaussian distributions. The dominant population, accounting for 84.7% of the
339 dataset, has a mean pCO₂ of 484 μatm, and is interpreted as representative of ambient background
340 seawater (Fig. 5); while populations 2 and 3, accounting for 12.9% and 2.4% of the dataset, are
341 interpreted as reflecting above background pCO₂ levels in far-from-hydrothermal seep (e.g., at the
342 margin of bubbling areas) and close-to-hydrothermal seep (at and above bubbling areas) environments,
343 respectively (see Fig. 5). The mean pCO₂ of these two sub-populations are 753 and 2770 μatm,
344 respectively. Hence, subtracting a background seawater TDIC of 2297 μmol/kg (corresponding to a
345 pCO₂ of a 484 μatm), the TDIC calculated above (2280-2830 μmol/kg) convert into Excess (above
346 background seawater) Dissolved Carbon (TDIC_{excess}) contents of 0 to 534 μmol/kg. We interpolate and
347 integrate such TDIC_{excess} results across the whole cross-sectional area AA' of Fig. 3b, using a Sequential
348 Gaussian Simulation⁷⁹, to obtain an integrated Excess Carbon content of 2.7±0.03 mol·m²/kg.

349 Finally, we estimate the advective transport of Excess Carbon across transect AA' using seawater
350 current velocities derived from an ocean circulation model for the Baia–Bagnoli area (see Methods and
351 Fig. S6). The vertical profile of the normal velocity $vn(z)$ and the associated transport $Q(z)$ reveal a
352 clear two-layer circulation (Fig. S6a), with net inflow in the upper water column and compensating
353 outflow at depth. Considering that the maximum depth of section AA' is ~15 m, we here rely on velocity
354 results for the upper layer, in which the mean current velocity is estimated at -0.0041 ± 0.0121 m/s
355 (Fig. S6a). The negative mean value obtained, and its large associated standard deviation, concur to
356 indicate that a weak (but temporally variable) water inflow toward the gulf prevails. For comparison,
357 the mean velocity at surface (0 m) returned by the circulation model is a factor circa 2 higher (-0.0076
358 ± 0.021 m/s; Fig. S6b). Combining these velocities with the integrated Excess Carbon content ($2.7 \pm$
359 0.03 mol·m²/kg), the Excess Carbon flux is estimated at 11.3 mol/s (≈ 43.1 tons CO₂/day) using the
360 upper-layer mean velocity (0.0041 m/s), and 21.1 mol/s (≈ 80.4 tons CO₂/day) when the model-derived
361 surface velocity (0.0076 m/s) is considered instead. It is important to keep in mind that the relatively
362 low mean current velocities inferred above are partially a result of the coexistence of negative (landward
363 current flow) and positive (seaward current flow) values. If we consider for example the upper range
364 ranges of landward (-0.0286 m/s) and seaward ($+0.0134$ m/s) surface velocities, the excess CO₂ flux
365 would increase to 302 tons and 141 tons CO₂/day, respectively. Similar results (fluxes ranging between
366 84 and 170 tons CO₂/day) would be obtained using the full velocity range across the water column in
367 the 0-15m vertical profile (Fig. S6a). Overall, these results suggest that the advective CO₂ transport
368 across the section falls within the 43–302 tons CO₂/day range.

369 Our calculations above have large associated uncertainties. For example, the ocean circulation model
370 simulations used for our flux estimations robustly reproduce the overall circulation pattern within the
371 basin and provide, to our knowledge, the most comprehensive state-of-the-art characterization currently
372 available for this area. However, although the model horizontal resolution is sufficient to resolve the

373 main coastal circulation patterns, some submesoscale processes and nearshore shear may remain
374 partially unresolved. This limitation may introduce moderate biases in pointwise velocity estimates,
375 likely leading to a slight underestimation of local velocity extremes. Nevertheless, the maximum
376 simulated velocities (of about 0.1 m/s) are consistent with the range reported from ADCP (Acoustic
377 Doppler Current Profiler) and drifter observations in the Bagnoli–Coroglio Bay within the Gulf of
378 Pozzuoli⁸⁰, providing confidence in our results, and supporting the overall reliability of the model in
379 representing the magnitude and structure of the circulation within the Gulf of Pozzuoli.

380 We also caution that the quality of our calculations is limited by the small density of our observation
381 points (see Fig. 1) that renders the extrapolation and integration procedure highly inaccurate. We cannot
382 exclude, or even consider likely, that other substantial emission zones exist in the bay that may have
383 been missed by our sparse observations. Conversely, we cannot exclude that our sampling grid is
384 somewhat biased by oversampling emission zones with visible surface manifestations (e.g., known
385 areas of bubble emission). We hence advise our calculations are only to be viewed as an order of
386 magnitude assessment of the CO₂ flux.

387 With these limitations in mind, we note our inferred CO₂ flux range (43–302 tons/day) for the
388 submerged CFC sector corresponds to circa 1.4–10% of the CO₂ flux from the Solfatara Diffuse
389 Degassing Structure (CO₂ flux of ~3000 tons/day), the main degassing structure in the caldera
390 onshore^{45,31,28,81}. This strongly exhaling area, centered on the Solfatara tuff crater^{67,68} (Fig. 1, 4), has been
391 the primary target of geochemical monitoring so far²⁸. While emphasising that additional measurements
392 - with a denser observation grid - are required for a refined CO₂ flux assessment, our results here suggest
393 that while CO₂ degassing today predominantly occurs in the inland CFC sector, the caldera offshore is
394 a non-trivial source of CO₂. More effective integration of the submerged caldera sector in ongoing
395 geochemical volcano monitoring efforts is hence recommended.

396 **Impact of CO₂ degassing on microbial diversity.** The widespread and structurally controlled release
397 of CO₂-rich fluids documented across the submerged sector of the Campi Flegrei caldera (Fig. 2) exerts
398 a first-order control on the chemical and biological organization of the overlying microbial diversity,
399 with potential impacts on the entire gulf marine ecosystem. Rather than representing a perturbation,
400 CO₂ degassing generates sharp and persistent geochemical gradients in the seawater column (Fig. 3),
401 locally modifying pH, dissolved inorganic carbon availability, redox conditions, and, as expected from
402 studies from other shallow water hydrothermal vents worldwide^{82,83}, influencing the delivery of reduced
403 compounds and trace metals. Together, these factors define a mosaic of highly heterogeneous
404 microenvironments that act as ecosystem engineers, selecting for specific metabolic strategies and
405 structuring microbial communities over spatial scales ranging from centimeters to hundreds of
406 meters^{10,13,84}.

407 Our results on the microbial diversity show that active degassing areas host dense and functionally
408 distinct microbial assemblages dominated by chemolithoautotrophic and mixotrophic lineages,
409 particularly within vent fluids and associated biofilms (Fig. 6). These communities are enriched in
410 metabolic pathways linked to carbon fixation and sulfur cycling that are characteristic of hydrothermal
411 systems (Fig. 8), including the reductive tricarboxylic acid cycle and the Wood–Ljungdahl pathway,
412 alongside more broadly distributed pathways such as the Calvin–Benson–Bassham cycle. Importantly,
413 these functional signatures are not confined to the immediate vent environment. Instead, genes and taxa
414 typically associated with hydrothermal activity are detectable—albeit at much lower relative abundance
415—in the surrounding seawater column, with systematic gradients linked to depth, distance from the
416 vent, and local pH (Fig. 6, 7 and 8). This pattern is consistent with dilution and mixing of hydrothermal
417 plumes into ambient seawater, indicating that CO₂ degassing sites act as conduits exporting subsurface-
418 derived microorganisms and metabolic potential into the coastal ocean^{85,86}. This is in line with early
419 reports of subsurface microorganisms venting from deep sea hydrothermal vents⁸⁷, and later studies
420 using venting sites as windows to study the subsurface biosphere^{88–90}.

421 While our data do not allow quantification of *in situ* carbon fixation rates, the observed functional
422 imprint strongly suggests that CO₂ degassing may locally enhance microbial primary productivity by
423 increasing carbon availability and sustaining redox disequilibria. Hydrothermal fluids at Campi Flegrei
424 are known to transport not only CO₂ but also reduced sulfur species, iron, and other trace metals^{33,44},
425 which can further stimulate chemolithotrophic metabolism and relieve micronutrient limitations in
426 coastal waters. Similar coupling between volcanic CO₂ release, metal enrichment, and elevated
427 microbial productivity has been reported in other shallow-water hydrothermal systems and natural CO₂
428 seeps^{91,92}, supporting the interpretation that degassing-driven inputs can subsidize local biogeochemical
429 cycling rather than simply imposing stress. In this context, CO₂ acts not in isolation but as part of a
430 broader geochemical package that reshapes the metabolic landscape of the ecosystem.

431 At the same time, the ecological consequences of CO₂ degassing are inherently scale- and taxa-
432 dependent (Fig. 6 and 7), and might affect meiofauna and macroorganism differently^{93,94}. The extreme
433 pH conditions measured near active vents are expected to impose strong physiological constraints on
434 many eukaryotic organisms and higher trophic levels, consistent with observations from other acidified
435 volcanic settings^{95,96}. These stress effects are likely to result in localized exclusion or reduced
436 performance of sensitive taxa^{97,98}. However, for microorganisms, the conditions documented here
437 impose a forcing on the community composition selecting for taxa capable to exploit hydrothermal
438 conditions, rather than constituting a stressor for mesophilic, background condition adapted
439 communities. Our data indicate that microbial communities respond primarily through compositional
440 and functional reorganization, with shifts toward metabolisms capable of exploiting elevated CO₂
441 concentrations, reduced compounds, and metal availability. Thus, CO₂ degassing simultaneously

442 generates zones of ecological stress for some components of the ecosystem, mainly meiofauna and
443 macrofauna as extensively documented in the literature^{99,100} and zones of enhanced opportunity for
444 microorganisms^{8,15}, reinforcing the view of volcanic degassing as a driver of ecological diversity.
445 Additionally, microorganisms present on the seafloor and in the shallow crust and active within the
446 hydrothermal system might contribute to alter the composition and quantity of volatiles recycled to the
447 surface as previously reported for other volcanic degassing areas¹⁶⁻¹⁸, possibly further limiting our
448 estimates of CO₂ released to the water column and atmosphere.

449 More broadly, the Campi Flegrei exemplifies how submerged calderas can function as persistent
450 interfaces between the deep Earth and the ocean, coupling magmatic degassing, subsurface microbial
451 ecosystems, and coastal marine biogeochemistry. The export of hydrothermal microorganisms and
452 functions into the water column implies a previously underappreciated pathway by which the subsurface
453 biosphere can influence nearshore ecosystems. In addition, the net export of CO₂, nutrients and trace
454 metals from the venting area to the surrounding environments as exemplified by the detected pCO₂
455 plumes has the potential to fertilize primary productivity with positive cascading effects on the local
456 coastal diversity^{101,102}. Trace metals released from water rock interactions might be particularly
457 important in controlling biogeochemical cycles in the surrounding areas^{103,104} by altering the available
458 metabolic pathways available to microorganisms¹⁰⁵. Given the abundance of partially or fully
459 submerged calderas worldwide, similar processes are likely to operate far beyond Campi Flegrei, with
460 implications for regional carbon cycling, nutrient dynamics, and ecosystem structure that are not
461 captured by studies focused solely on subaerial degassing or open-ocean hydrothermal systems.

462 Taken together, our results show that submarine CO₂ degassing in underwater calderas is not merely a
463 geochemical expression of volcanic unrest, but a dynamic ecological force capable of reorganizing
464 microbial community structure and function at large scales influencing surrounding ecosystems.
465 Integrating biological observations with high-resolution degassing surveys reveals submerged calderas
466 as active biogeochemical reactors, where Earth processes and marine ecosystems are tightly and
467 continuously coupled.

468 **Methods**

469 **pCO₂ underwater sensors.** The seawater results from the Pozzuoli Bay were obtained during nine
470 different surveys performed from June 2023 to November 2025 using appropriate underwater CO₂
471 sensors: the HydroCTM CONTROS System and Solutions (Kiel, Germany) and the CO₂-ProTM
472 OCEANUS CV Submersible pCO₂ Sensor. Both sensors are compact submersible systems developed
473 for *in situ* determination of the partial pressure of carbon dioxide (pCO₂) in seawater. Their analytical
474 principle is based on nondispersive infrared (NDIR) spectrometry, whereby dissolved CO₂ permeates a
475 gas-permeable membrane into an internal equilibration volume^{33,106}. Continuous flushing of the

476 membrane surface is ensured by an external pump, which both supplies seawater to the sensor head and
477 minimizes boundary layer effects that could otherwise impair diffusion rate¹⁰⁶. Data acquisition was
478 carried out at a frequency of 1 Hz and transmitted in real time to a computer via a 50 m communication
479 cable. The HydroC™ CONTROS System was calibrated across a temperature range of 13–30 °C, with
480 a working pCO₂ interval between 100 and 6000 µatm (10–600 Pa). Measurement precision is better
481 than 1 µatm (0.1 Pa). To mitigate instrumental drift, the sensor was operated with regular automated
482 sequences consisting of 120s zeroing intervals and 480s flushing phases, the results of which were
483 applied during subsequent calibration correction¹⁰⁶. The characteristic response time (T₆₃) was
484 determined to be 20 ± 2s near the surface and approximately 30s at ~20 m depth. CO2-Pro™
485 OCEANUS CV Submersible pCO₂ Sensor was designed to accommodate a broader concentration
486 spectrum—from nominal ranges (e.g., 0–600 ppm) up to 10 000 ppm and was used in the areas where
487 the pCO₂ exceeded the 2000 µatm. The instrument's measurement is underpinned by high accuracy
488 (±0.5%) and a fine resolution of 0.01 ppm, enhanced through automatic zero compensation and factory
489 calibration using WMO-traceable (World Meteorological Organization) standards gases for long-term
490 stability. Equilibration time (t₆₃) was approximately 50 seconds with a pumped head at 20 °C, yielding
491 rapid sensor responsiveness under flow conditions.

492 Seawater vertical profiles were obtained along a relatively regular grid with a spacing between 100-
493 200m, particularly, in the areas where pCO₂ was relatively low (close to the normal seawater value
494 ~400-500 µatm⁴³). During each survey, vertical profiles, far from the hydrothermal areas, were acquired
495 in order to check the local seawater background. The spacing was reduced (10–50m) near hydrothermal
496 vents (e.g., Secca delle Fumose, Baia Castle beach, Academia-Bagnoli-Nisida coast and Capo Miseno
497 area). A total of 273 geochemical profiles were performed covering an area of approximately ~19 km².
498 At each site, the CO₂ sensor was lowered from a small boat down to the seafloor, and then raised up to
499 the seawater surface at a rate of 20cm/s. While performing the measurements along each vertical profile,
500 and at each 4-5m step, the CO₂ sensor was stopped and maintained at a constant depth for a few minutes
501 (2-3), in order to allow the sensors to fully equilibrate. After equilibration and based on the depth of the
502 seafloor, we measured the pCO₂ at 1 (0-4m), 2 (4-8m) or 3 (8-15m) different depths for each profile.
503 Finally, given the irregularity of the seafloor in the areas of Secca delle Fumose and Capo Miseno, the
504 vertical profiling was assisted by a scuba diver to ensure that the sensors measured the pCO₂ near-vent
505 plumes.

506 The final maps were developed using the maximum pCO₂ registered at each depth. Spatial interpolation
507 and mapping (Figs. 2-3) were performed in Surfer® (from Golden Software, LLC;
508 www.goldensoftware.com¹⁰⁷) using ordinary kriging, a geostatistical interpolation method that
509 explicitly models spatial autocorrelation; empirical semivariograms were computed and fitted with a
510 theoretical model, which was then used to estimate values at unsampled locations and generate

511 continuous spatial maps. Semivariogram parameters (nugget, sill, and range) were estimated by
 512 weighted least squares, and kriging predictions were computed on a regular grid covering the study
 513 area.

514 Seawater pCO₂ profiles (Fig. 3) were generated using the Data-Interpolating Variational Analysis
 515 (DIVA) gridding method implemented in Ocean Data View (ODV)¹⁰⁸. DIVA variational interpolation
 516 technique produces spatially consistent fields by minimizing a cost function constrained by data misfit
 517 and smoothness; in this study, gridded pCO₂ values were extracted along the vertical dimension to
 518 obtain continuous profiles. DIVA gridding was applied using default ODV parameters, with correlation
 519 length scales and signal-to-noise ratios optimized internally by the software. The resulting three-
 520 dimensional gridded fields were used to derive vertical seawater pCO₂ profiles.

521 **Modelling sea water current in the Pozzuoli bay.** Surface and subsurface currents in the Gulf of
 522 Pozzuoli were estimated using a high-resolution regional ocean hindcasts of the Gulf of Naples for 2016
 523 based on the Coastal Regional Ocean Model (CROM)¹⁰⁹. The model domain (40°–41.43° N, 13°–15.36°
 524 E) was implemented on a regular grid with horizontal resolution of ≈ 600 m and 40 terrain-following
 525 sigma levels, with the bathymetry derived from the 30" GEBCO dataset. Initial and open-boundary
 526 conditions consistent with the basin-scale transports were guaranteed through the one-way nesting¹¹⁰
 527 within the Mediterranean NEMO-OPA circulation model¹¹¹. Surface boundary conditions, i.e.
 528 momentum from wind, heat and freshwater fluxes, were computed from the 5-km resolution
 529 SKIRON/Eta atmospheric model at a temporal step of 6 hours. The simulation consisted of twelve
 530 monthly hindcasts for the year analysed, with monthly re-initialisation using satellite-derived sea
 531 surface temperature (SST) provided by the Copernicus Marine Environment Monitoring Service
 532 (CMEMS) to constrain the surface heat fluxes¹¹². Every simulation was initialized one week prior to the
 533 target month to allow model adjustment, so that the outputs of the first six days were discarded to
 534 remove potential initialisation transients associated with the spin-up phase.

535 Volume transport across the AA' transect was computed by integrating the velocity component normal
 536 to the section across the full water column:

$$537 \quad Q = \sum_{k,s} v_n(k, s) A(k, s) \quad (1)$$

538

539

540 where v_n is the velocity component normal to the section and $A(k,s)$ is the cross-sectional area
 541 associated with each grid segment and sigma layer. Land points were excluded from the calculation.

542 **Biological samples collection.** Biofilm samples and hydrothermal fluids were collected by divers at
543 both hot, focused and warm, diffuse venting sites, minimizing disturbance of the sampled matrices.
544 Seawater column samples were collected from the vessel using Niskin bottles following a depth-
545 stratified approach: (i) a single depth when seafloor depth was <6 m below seafloor (bsf); (ii) two depths
546 when seafloor depth ranged between 6 and 8 m bsf; and (iii) three depths when seafloor depth exceeded
547 8 m bsf. The maximum sampling depth was 15 m bsf. In total, 27 samples were collected for biological
548 analyses, including biofilms (n = 5), hydrothermal fluid (n = 1), and seawater (n = 21). Samples
549 SF1D13_B1 and SF1D13_B2 correspond to orange biofilms collected directly at the SdF emission site
550 in June and July 2023, respectively. Samples SF2A_B1, SF2B_B1, and SF2C_B1 correspond to white
551 biofilms collected from colder seafloor areas characterized by diffuse hydrothermal venting. Sample
552 SF1D13_F2 represents the hydrothermal fluid collected at the SdF emission site. Seawater samples
553 (TxDy) were collected from the water column surrounding SdF at distinct transect sites and depths (x
554 = transect site; y = depth; Fig. 2b). Hydrothermal fluid and seawater samples were filtered in the field
555 through 0.22 µm Sterivex filter membranes for biomass concentration. Biofilm samples from the SdF
556 vent and surrounding venting seafloor were collected in sterile 50-mL centrifuge tubes. In the
557 laboratory, Sterivex filters and biofilm samples were stored at -20 °C until DNA extraction. Physico-
558 chemical parameters of fluids were measured *in situ* using a multiparametric probe (HANNA,
559 HI98494). Fluid aliquots were filtered through 0.22 µm membranes for immediate determination of
560 total alkalinity.

561

562 **DNA extraction and shotgun metagenomics.** DNA was extracted from Sterivex filters and biofilm
563 samples using a modified phenol–chloroform extraction protocol optimized for shallow-water
564 hydrothermal vents and seawater samples¹⁰. DNA concentration was quantified fluorometrically using
565 a Qubit (Thermo Fisher Scientific). Shotgun metagenomic sequencing was performed at Novogene with
566 the NExtSeq Illumina platform (PE 2x150) with a sequencing depth between 15 and 30 million reads
567 per sample.

568 **Bioinformatic and statistical analysis.** Raw shotgun metagenomic reads were processed using the
569 GeoMosaic pipeline (Giovannelli Lab, <https://giovannellilab.github.io/Geomosaic/index.html>), which
570 performs integrated quality control, taxonomic profiling, functional annotation, and metagenome
571 assembly. Sequencing reads were first quality-filtered and trimmed within GeoMosaic using Fastp,
572 Fastqc and Reads Count; then, high-quality reads were functionally annotated using Kaiju, employing
573 a protein-level classification strategy suitable for hydrothermal vent and marine metagenomes^{113–115}.
574 Functional annotation of the same reads was carried out using FunProfiler together with a custom
575 database of KEGG orthologues (KOs) associated with biogeochemical cycles¹¹⁶. Diversity analyses
576 were conducted in R using the *phyloseq* package within a Jupyter notebook instance¹¹⁷. Beta diversity

577 was evaluated using weighted Jaccard distance metrics, as implemented in the vegan package¹¹⁸. All
578 statistical analyses and visualizations were performed in R v4.5.1 (R Core Team, 2021) using
579 *ggplot2*¹¹⁹.

580 Data availability

581 The datasets generated and/or analyzed during the current study are available in the Supplementary
582 Tables provided with this manuscript. All sequencing data generated in this study, including 16S rRNA
583 gene amplicons and shotgun metagenomes, have been deposited in the European Nucleotide Archive
584 (ENA) under the project accession number PRJEB87249 under the Umbrella Project CoEvolve
585 PRJEB55081. Sequences from the Barosa et al. (2025) study are available under access number
586 PRJEB67762 under the same Umbrella Project. The bioinformatic code used for data processing and
587 analysis, together with associated metadata files, is publicly available in the GitHub repository
588 https://github.com/giovannellilab/Sandoval-Velasquez_Campi_Flegrei_degassing/tree/main and
589 archived on Zenodo under DOI 10.5281/zenodo.18800558.

590 References

- 591 1. Branney, M. & Acocella, V. Calderas. in *The Encyclopedia of Volcanoes* 299–315 (Elsevier, 2015).
592 doi:10.1016/B978-0-12-385938-9.00016-X.
- 593 2. Rouwet, D. *Volcanic Lakes*. (Springer Berlin / Heidelberg, Berlin, Heidelberg, 2015).
- 594 3. Fouquet, Y. et al. Volcanic and hydrothermal processes in submarine calderas: The Kulo Lasi
595 example (SW Pacific). *Ore Geol. Rev.* **99**, 314–343 (2018).
- 596 4. Enochs, I. C. et al. Shift from coral to macroalgae dominance on a volcanically acidified reef. *Nat.*
597 *Clim. Change* **5**, 1083–1088 (2015).
- 598 5. Tarasov, V. G. et al. Effect of shallow-water hydrothermal venting on the biota of Matupi Harbour
599 (Rabaul Caldera, New Britain Island, Papua New Guinea). *Cont. Shelf Res.* **19**, 79–116 (1999).
- 600 6. Aiuppa, A. et al. Volcanic CO₂ seep geochemistry and use in understanding ocean acidification.
601 *Biogeochemistry* **152**, 93–115 (2021).
- 602 7. Agostini, S. et al. Geochemistry of two shallow CO₂ seeps in Shikine Island (Japan) and their
603 potential for ocean acidification research. *Reg. Stud. Mar. Sci.* **2**, 45–53 (2015).
- 604 8. Price, R. E. & Giovannelli, D. A Review of the Geochemistry and Microbiology of Marine
605 Shallow-Water Hydrothermal Vents. in *Reference Module in Earth Systems and Environmental*
606 *Sciences* B9780124095489095233 (Elsevier, 2017). doi:10.1016/B978-0-12-409548-9.09523-3.
- 607 9. Sievert, S. & Vetriani, C. Chemoautotrophy at Deep-Sea Vents: Past, Present, and Future.
608 *Oceanography* **25**, 218–233 (2012).
- 609 10. Giovannelli, D., d’Errico, G., Manini, E., Yakimov, M. & Vetriani, C. Diversity and phylogenetic
610 analyses of bacteria from a shallow-water hydrothermal vent in Milos island (Greece). *Front.*
611 *Microbiol.* **4**, (2013).
- 612 11. Barosa, B. et al. Mapping the microbial diversity associated with different geochemical regimes in
613 the shallow-water hydrothermal vents of the Aeolian archipelago, Italy. *Front. Microbiol.* **14**,
614 1134114 (2023).
- 615 12. Barosa, B. et al. The microbiology and geochemistry of the shallow-water hydrothermal vents of
616 the Gulf of Naples, Italy. Preprint at <https://doi.org/10.1101/2025.10.02.679971> (2025).
- 617 13. Yücel, M. et al. Eco-geochemical dynamics of a shallow-water hydrothermal vent system at Milos
618 Island, Aegean Sea (Eastern Mediterranean). *Chem. Geol.* **356**, 11–20 (2013).

- 619 14. Gomez-Saez, G. V. *et al.* Relative Importance of Chemoautotrophy for Primary Production in a
620 Light Exposed Marine Shallow Hydrothermal System. *Front. Microbiol.* **8**, 702 (2017).
- 621 15. Giovannelli, D. & Price, R. E. Marine Shallow-Water Hydrothermal Vents: Microbiology. in
622 *Encyclopedia of Ocean Sciences* 353–363 (Elsevier, 2019). doi:10.1016/B978-0-12-409548-
623 9.11250-3.
- 624 16. Barry, P. H. *et al.* Forearc carbon sink reduces long-term volatile recycling into the mantle.
625 *Nature* **568**, 487–492 (2019).
- 626 17. Fullerton, K. M. *et al.* Effect of tectonic processes on biosphere-geosphere feedbacks across a
627 convergent margin. *Nat. Geosci.* **14**, 301–306 (2021).
- 628 18. Giovannelli, D. *et al.* Subsurface life can modify volatile cycling on a planetary scale. *Mem. Della*
629 *Soc. Astron. Ital.* **92**, 60–63 (2021).
- 630 19. Del Gaudio, C., Aquino, I., Ricciardi, G. P., Ricco, C. & Scandone, R. Unrest episodes at Campi
631 Flegrei: A reconstruction of vertical ground movements during 1905–2009. *J. Volcanol. Geotherm.*
632 *Res.* **195**, 48–56 (2010).
- 633 20. Scarpa, R. *et al.* Historic Unrest of the Campi Flegrei Caldera, Italy. in *Campi Flegrei* (eds Orsi,
634 G., D’Antonio, M. & Civetta, L.) 257–282 (Springer Berlin Heidelberg, Berlin, Heidelberg, 2022).
635 doi:10.1007/978-3-642-37060-1_10.
- 636 21. Kilburn, C. R. J., Carlino, S., Danesi, S. & Pino, N. A. Potential for rupture before eruption at
637 Campi Flegrei caldera, Southern Italy. *Commun. Earth Environ.* **4**, 190 (2023).
- 638 22. Astort, A. *et al.* Tracking the 2007–2023 magma-driven unrest at Campi Flegrei caldera (Italy).
639 *Commun. Earth Environ.* **5**, 506 (2024).
- 640 23. Orsi, G. *et al.* Short-term ground deformations and seismicity in the resurgent Campi Flegrei
641 caldera (Italy): an example of active block-resurgence in a densely populated area. *J. Volcanol.*
642 *Geotherm. Res.* **91**, 415–451 (1999).
- 643 24. Tramelli, A., Convertito, V. & Godano, C. b value enlightens different rheological behaviour in
644 Campi Flegrei caldera. *Commun. Earth Environ.* **5**, 275 (2024).
- 645 25. Scotto Di Uccio, F. *et al.* Delineation and Fine-Scale Structure of Fault Zones Activated During
646 the 2014–2024 Unrest at the Campi Flegrei Caldera (Southern Italy) From High-Precision
647 Earthquake Locations. *Geophys. Res. Lett.* **51**, e2023GL107680 (2024).
- 648 26. De Landro, G. *et al.* 3D structure and dynamics of Campi Flegrei enhance multi-hazard assessment.
649 *Nat. Commun.* **16**, 4814 (2025).
- 650 27. Tan, X. *et al.* A clearer view of the current phase of unrest at Campi Flegrei caldera. *Science* **390**,
651 70–75 (2025).
- 652 28. Chiodini, G. *et al.* The Hydrothermal System of the Campi Flegrei Caldera, Italy. in *Campi*
653 *Flegrei* (eds Orsi, G., D’Antonio, M. & Civetta, L.) 239–255 (Springer Berlin Heidelberg, Berlin,
654 Heidelberg, 2022). doi:10.1007/978-3-642-37060-1_9.
- 655 29. Chiodini, G. *et al.* Magmas near the critical degassing pressure drive volcanic unrest towards a
656 critical state. *Nat. Commun.* **7**, 13712 (2016).
- 657 30. Caliro, S. *et al.* Escalation of caldera unrest indicated by increasing emission of isotopically light
658 sulfur. *Nat. Geosci.* **18**, 167–174 (2025a).
- 659 31. Cardellini, C. *et al.* Monitoring diffuse volcanic degassing during volcanic unrests: the case of
660 Campi Flegrei (Italy). *Sci. Rep.* **7**, 6757 (2017).
- 661 32. Caliro, S. *et al.* Chemical and isotopic characterization of groundwater and thermal waters from the
662 Campi Flegrei caldera (southern Italy). *J. Volcanol. Geotherm. Res.* **460**, 108280 (2025b).
- 663 33. Di Napoli, R. *et al.* Hydrothermal fluid venting in the offshore sector of C ampi F legrei caldera: A
664 geochemical, geophysical, and volcanological study. *Geochem. Geophys. Geosystems* **17**, 4153–
665 4178 (2016).
- 666 34. Passaro, S. *et al.* Multi-resolution morpho-bathymetric survey results at the Pozzuoli–Baia
667 underwater archaeological site (Naples, Italy). *J. Archaeol. Sci.* **40**, 1268–1278 (2013).
- 668 35. Natale, J. *et al.* Integrated on-land-offshore stratigraphy of the Campi Flegrei caldera: New insights
669 into the volcano-tectonic evolution in the last 15 kyr. *Basin Res.* **34**, 855–882 (2022a).
- 670 36. Natale, J. *et al.* Fault systems in the offshore sector of the Campi Flegrei caldera (southern Italy):
671 Implications for nested caldera structure, resurgent dome, and volcano-tectonic evolution. *J. Struct.*
672 *Geol.* **163**, 104723 (2022b).

- 673 37. Steinmann, L., Spiess, V. & Sacchi, M. The Campi Flegrei caldera (Italy): Formation and
674 evolution in interplay with sea-level variations since the Campanian Ignimbrite eruption at 39 ka. *J.*
675 *Volcanol. Geotherm. Res.* **327**, 361–374 (2016).
- 676 38. Steinmann, L., Spiess, V. & Sacchi, M. Post-collapse evolution of a coastal caldera system:
677 Insights from a 3D multichannel seismic survey from the Campi Flegrei caldera (Italy). *J. Volcanol.*
678 *Geotherm. Res.* **349**, 83–98 (2018).
- 679 39. Sacchi, M. *et al.* The Neapolitan Yellow Tuff caldera offshore the Campi Flegrei: Stratal
680 architecture and kinematic reconstruction during the last 15ky. *Mar. Geol.* **354**, 15–33 (2014).
- 681 40. Aucelli, P., Cinque, A., Mattei, G., Pappone, G. & Rizzo, A. Studying relative sea level change
682 and correlative adaptation of coastal structures on submerged Roman time ruins nearby Naples
683 (southern Italy). *Quat. Int.* **501**, 328–348 (2019).
- 684 41. Aucelli, P. P. C. *et al.* Multi-proxy analysis of relative sea-level and paleoshoreline changes during
685 the last 2300 years in the Campi Flegrei caldera, Southern Italy. *Quat. Int.* **602**, 110–130 (2021).
- 686 42. Baldrighi, E. *et al.* Meiofaunal communities and nematode diversity characterizing the Secca delle
687 Fumose shallow vent area (Gulf of Naples, Italy). *PeerJ* **8**, e9058 (2020).
- 688 43. Berry. The impact of human CO₂ on atmospheric CO₂. *Sci. Clim. Change Int. J. Sci. Philos.* **36**
689 (2022) doi:10.53234/scc202112/212.
- 690 44. Vaselli, O., Franco, T., Dario, T., Robert, P. J. & Antonio, C. Submarine and Inland Gas
691 Discharges from the Campi Flegrei (Southern Italy) and the Pozzuoli Bay: Geochemical Clues for
692 a Common Hydrothermal-Magmatic Source. *Procedia Earth Planet. Sci.* **4**, 57–73 (2011).
- 693 45. Buono, G., Caliro, S., Paonita, A., Pappalardo, L. & Chiodini, G. Discriminating carbon dioxide
694 sources during volcanic unrest: The case of Campi Flegrei caldera (Italy). *Geology* **51**, 397–401
695 (2023).
- 696 46. Campbell, B. J., Engel, A. S., Porter, M. L. & Takai, K. The versatile ϵ -proteobacteria: key players
697 in sulphidic habitats. *Nat. Rev. Microbiol.* **4**, 458–468 (2006).
- 698 47. Giovannelli, D. *et al.* Insight into the evolution of microbial metabolism from the deep-branching
699 bacterium, *Thermovibrio ammonificans*. *eLife* **6**, e18990 (2017).
- 700 48. Hügler, M. & Sievert, S. M. Beyond the Calvin Cycle: Autotrophic Carbon Fixation in the Ocean.
701 *Annu. Rev. Mar. Sci.* **3**, 261–289 (2011).
- 702 49. Pontiller, B. *et al.* Rapid bacterioplankton transcription cascades regulate organic matter utilization
703 during phytoplankton bloom progression in a coastal upwelling system. *ISME J.* **16**, 2360–2372
704 (2022).
- 705 50. Buchan, A., LeClerc, G. R., Gulvik, C. A. & González, J. M. Master recyclers: features and
706 functions of bacteria associated with phytoplankton blooms. *Nat. Rev. Microbiol.* **12**, 686–698
707 (2014).
- 708 51. Newhall, C. G. & Dzurisin, D. *Historical Unrest at Large Calderas of the World.*
709 <https://pubs.usgs.gov/publication/b1855> (1988) doi:10.3133/b1855.
- 710 52. Acocella, V., Di Lorenzo, R., Newhall, C. & Scandone, R. An overview of recent (1988 to 2014)
711 caldera unrest: Knowledge and perspectives. *Rev. Geophys.* **53**, 896–955 (2015).
- 712 53. Lowenstern, J. B., Smith, R. B. & Hill, D. P. Monitoring super-volcanoes: geophysical and
713 geochemical signals at Yellowstone and other large caldera systems. *Philos. Trans. R. Soc. Math.*
714 *Phys. Eng. Sci.* **364**, 2055–2072 (2006).
- 715 54. Lowenstern, J. B. & Hurwitz, S. Monitoring a Supervolcano in Repose: Heat and Volatile Flux at
716 the Yellowstone Caldera. *Elements* **4**, 35–40 (2008).
- 717 55. Aiuppa, A., Caudron, C., Chiodini, G., Ingebritsen, S. & Viveiros, F. Geochemical Monitoring of
718 Volcanic Fluids in the Twenty-First Century. in *Modern Volcano Monitoring* (eds Spica, Z. &
719 Caudron, C.) 209–260 (Springer Nature Switzerland, Cham, 2025a). doi:10.1007/978-3-031-
720 86841-2_8.
- 721 56. Aiuppa, A. *et al.* Geochemical volcano monitoring. Preprint at <https://doi.org/10.31223/X5TT6D>
722 (2025b).
- 723 57. Dura, A. *et al.* The Hydrothermal Vent Field at the Eastern Edge of the Hellenic Volcanic Arc: The
724 Avyssos Caldera (Nisyros). *Geosciences* **11**, 290 (2021).
- 725 58. Agostini, S. *et al.* Ocean acidification drives community shifts towards simplified non-calcified
726 habitats in a subtropical–temperate transition zone. *Sci. Rep.* **8**, 11354 (2018).

- 727 59. De Ronde, C. E. J. *et al.* Submarine hydrothermal activity along the mid-Kermadec Arc, New
728 Zealand: Large-scale effects on venting. *Geochem. Geophys. Geosystems* **8**, 2006GC001495
729 (2007).
- 730 60. Baker, E. T. *et al.* Hydrothermal activity and volcano distribution along the Mariana arc. *J.*
731 *Geophys. Res. Solid Earth* **113**, 2007JB005423 (2008).
- 732 61. Italiano, F. & Nuccio, P. M. Geochemical investigations of submarine volcanic exhalations to the
733 east of Panarea, Aeolian Islands, Italy. *J. Volcanol. Geotherm. Res.* **46**, 125–141 (1991).
- 734 62. Fischer, T. P. *et al.* The emissions of CO₂ and other volatiles from the world's subaerial
735 volcanoes. *Sci. Rep.* **9**, 18716 (2019).
- 736 63. Werner, C. *et al.* Carbon Dioxide Emissions from Subaerial Volcanic Regions: Two Decades in
737 Review. in *Deep Carbon* (eds Orcutt, B. N., Daniel, I. & Dasgupta, R.) 188–236 (Cambridge
738 University Press, 2019). doi:10.1017/9781108677950.008.
- 739 64. Fischer, T. P. & Aiuppa, A. AGU Centennial Grand Challenge: Volcanoes and Deep Carbon
740 Global CO₂ Emissions From Subaerial Volcanism—Recent Progress and Future Challenges.
741 *Geochem. Geophys. Geosystems* **21**, e2019GC008690 (2020).
- 742 65. Bastianoni, A. *et al.* The missing carbon budget puzzle piece: shallow-water hydrothermal vents
743 contribution to global CO₂ fluxes. Preprint at <https://doi.org/10.31223/X52W81> (2022).
- 744 66. Marini, L., Principe, C. & Lelli, M. *The Solfatara Magmatic-Hydrothermal System: Geochemistry,*
745 *Geothermometry and Geobarometry of Fumarolic Fluids.* (Springer International Publishing,
746 Cham, 2022). doi:10.1007/978-3-030-98471-7.
- 747 67. Di Vito, M. A. *et al.* Volcanism and deformation since 12,000 years at the Campi Flegrei caldera
748 (Italy). *J. Volcanol. Geotherm. Res.* **91**, 221–246 (1999).
- 749 68. Isaia, R. *et al.* Stratigraphy, structure, and volcano-tectonic evolution of Solfatara maar-diatreme
750 (Campi Flegrei, Italy). *Geol. Soc. Am. Bull.* **127**, 1485–1504 (2015).
- 751 69. Chiodini, G. *et al.* CO₂ degassing and energy release at Solfatara volcano, Campi Flegrei, Italy. *J.*
752 *Geophys. Res. Solid Earth* **106**, 16213–16221 (2001).
- 753 70. Giordano, G. *et al.* Birth and growth of a volcanotectonic fault during the current volcanic unrest
754 at Campi Flegrei caldera (Italy). *Commun. Earth Environ.* **6**, 839 (2025).
- 755 71. Giudicepietro, F. *et al.* Burst-like swarms in the Campi Flegrei caldera accelerating unrest from
756 2021 to 2024. *Nat. Commun.* **16**, 1548 (2025).
- 757 72. Gianfrotta, Piero. A. Harbor Structures of the Augustan Age in Italy. in *A retrospective after Two*
758 *Millennia. Atti del Convegno (Caesarea Maritima 1995)* 65–76 (Leiden-New York-Köln, 1996).
- 759 73. Scotto Di Uccio, F. *et al.* Delineation and Fine-Scale Structure of Fault Zones Activated During
760 the 2014–2024 Unrest at the Campi Flegrei Caldera (Southern Italy) From High-Precision
761 Earthquake Locations. *Geophys. Res. Lett.* **51**, e2023GL107680 (2024).
- 762 74. Tramelli, A., Giudicepietro, F., Ricciolino, P. & Chiodini, G. The seismicity of Campi Flegrei in
763 the contest of an evolving long term unrest. *Sci. Rep.* **12**, 2900 (2022).
- 764 75. Di Vito, M. A. *et al.* Magma transfer at Campi Flegrei caldera (Italy) before the 1538 AD
765 eruption. *Sci. Rep.* **6**, 32245 (2016).
- 766 76. Orsi, G., D'Antonio, M., Vita, S. D. & Gallo, G. The Neapolitan Yellow Tuff, a large-magnitude
767 trachytic phreatoplinian eruption: eruptive dynamics, magma withdrawal and caldera collapse. *J.*
768 *Volcanol. Geotherm. Res.* **53**, 275–287 (1992).
- 769 77. Orsi, G. Volcanic and Deformation History of the Campi Flegrei Volcanic Field, Italy. in *Campi*
770 *Flegrei* (eds Orsi, G., D'Antonio, M. & Civetta, L.) 1–53 (Springer Berlin Heidelberg, Berlin,
771 Heidelberg, 2022). doi:10.1007/978-3-642-37060-1_1.
- 772 78. Sharp, J. D. *et al.* CO₂SYsv3 for MATLAB. Zenodo <https://doi.org/10.5281/ZENODO.3950562>
773 (2023).
- 774 79. Deutsch, C. V. & Journel, A. G. *GSLIB: Geostatistical Software Library and User's Guide.* (Oxford
775 university press, New York, 1998).
- 776 80. Castagno, P. *et al.* Hydrographic and dynamical characterisation of the Bagnoli-Coroglio Bay (Gulf
777 of Naples, Tyrrhenian Sea). *Chem. Ecol.* **36**, 598–618 (2020).
- 778 81. Chiodini, G. *et al.* Long-term variations of the Campi Flegrei, Italy, volcanic system as revealed by
779 the monitoring of hydrothermal activity. *J. Geophys. Res. Solid Earth* **115**, 2008JB006258
780 (2010).

- 781 82. Pichler, T., Veizer, J. & Hall, G. E. M. The chemical composition of shallow-water hydrothermal
782 fluids in Tutum Bay, Ambitle Island, Papua New Guinea and their effect on ambient seawater. *Mar.*
783 *Chem.* **64**, 229–252 (1999).
- 784 83. Tarasov, V. G. Effects of Shallow-Water Hydrothermal Venting on Biological Communities of
785 Coastal Marine Ecosystems of the Western Pacific. in *Advances in Marine Biology* vol. 50 267–
786 421 (Elsevier, 2006).
- 787 84. Silva, A. C. P. *et al.* Tectonic setting shapes microbial biosynthetic potential across global
788 geothermal environments. Preprint at <https://doi.org/10.1101/2025.09.14.675129> (2025).
- 789 85. Dick, G. J. *et al.* The microbiology of deep-sea hydrothermal vent plumes: ecological and
790 biogeographic linkages to seafloor and water column habitats. *Front. Microbiol.* **4**, (2013).
- 791 86. Fortunato, C. S. & Huber, J. A. Coupled RNA-SIP and metatranscriptomics of active
792 chemolithoautotrophic communities at a deep-sea hydrothermal vent. *ISME J.* **10**, 1925–1938
793 (2016).
- 794 87. Deming, J. W. & Baross, J. A. Deep-sea smokers: Windows to a subsurface biosphere? *Geochim.*
795 *Cosmochim. Acta* **57**, 3219–3230 (1993).
- 796 88. Giovannelli, D. *et al.* Sampling across large-scale geological gradients to study geosphere–
797 biosphere interactions. *Front. Microbiol.* **13**, 998133 (2022).
- 798 89. Elkassas, S. M. *et al.* Metabolic and population profiles of active seafloor autotrophs in young
799 oceanic crust at deep-sea hydrothermal vents. *Appl. Environ. Microbiol.* **91**, e01868-25 (2025).
- 800 90. Reveillaud, J. *et al.* Seafloor microbial communities in hydrogen-rich vent fluids from
801 hydrothermal systems along the Mid-Cayman Rise. *Environ. Microbiol.* **18**, 1970–1987 (2016).
- 802 91. Pérez-Barrancos, C. *et al.* Shallow Hydrothermal Fluids Shape Microbial Dynamics at the Tagoro
803 Submarine Volcano (Canary Islands, Spain). *Environ. Microbiol.* **27**, e70052 (2025).
- 804 92. Gomez-Saez, G. V. *et al.* Hydrothermal iron - DOM interaction in a shallow-water hydrothermal
805 system off Dominica Island. 2 datasets Preprint at <https://doi.org/10.1594/PANGAEA.874263>
806 (2015).
- 807 93. Bellec, L. *et al.* Microbial Communities of the Shallow-Water Hydrothermal Vent Near Naples,
808 Italy, and Chemosynthetic Symbionts Associated With a Free-Living Marine Nematode. *Front.*
809 *Microbiol.* **11**, 2023 (2020).
- 810 94. Milazzo, M. *et al.* Biogenic habitat shifts under long-term ocean acidification show nonlinear
811 community responses and unbalanced functions of associated invertebrates. *Sci. Total Environ.*
812 **667**, 41–48 (2019).
- 813 95. Garilli, V. *et al.* Physiological advantages of dwarfing in surviving extinctions in high-CO₂ oceans.
814 *Nat. Clim. Change* **5**, 678–682 (2015).
- 815 96. Harvey, J., König, S. & Luguet, A. The effects of melt depletion and metasomatism on highly
816 siderophile and strongly chalcophile elements: S–Se–Te–Re–PGE systematics of peridotite
817 xenoliths from Kilbourne Hole, New Mexico. *Geochim. Cosmochim. Acta* **166**, 210–233 (2015).
- 818 97. Hall-Spencer, J. M. *et al.* Volcanic carbon dioxide vents show ecosystem effects of ocean
819 acidification. *Nature* **454**, 96–99 (2008).
- 820 98. Fabricius, K. E. *et al.* Losers and winners in coral reefs acclimatized to elevated carbon dioxide
821 concentrations. *Nat. Clim. Change* **1**, 165–169 (2011).
- 822 99. Martins Medeiros, I. P. & Souza, M. M. Acid times in physiology: A systematic review of the
823 effects of ocean acidification on calcifying invertebrates. *Environ. Res.* **231**, 116019 (2023).
- 824 100. Kroeker, K. J. *et al.* Impacts of ocean acidification on marine organisms: quantifying
825 sensitivities and interaction with warming. *Glob. Change Biol.* **19**, 1884–1896 (2013).
- 826 101. Zhang, Z. *et al.* Substantial trace metal input from the 2022 Hunga Tonga-Hunga Ha’apai
827 eruption into the South Pacific. *Nat. Commun.* **15**, 8986 (2024).
- 828 102. Caramanna, G., Sievert, S. M. & Bühring, S. I. Submarine Shallow-Water Fluid Emissions
829 and Their Geomicrobiological Imprint: A Global Overview. *Front. Mar. Sci.* **8**, 727199 (2021).
- 830 103. Hay Mele, B. *et al.* Oxidoreductases and metal cofactors in the functioning of the earth. *Essays*
831 *Biochem.* **67**, 653–670 (2023).
- 832 104. Giovannelli, D. Trace metal availability and the evolution of biogeochemistry. *Nat. Rev. Earth*
833 *Environ.* **4**, 597–598 (2023).

- 834 105. Ricciardelli, A. *et al.* Trace metals availability controls terminal electron acceptor utilization in
 835 *Escherichia coli*. 2025.01.08.631794 Preprint at <https://doi.org/10.1101/2025.01.08.631794>
 836 (2025).
- 837 106. Fietzek, P., Fiedler, B., Steinhoff, T. & Körtzinger, A. In situ Quality Assessment of a Novel
 838 Underwater pCO₂ Sensor Based on Membrane Equilibration and NDIR Spectrometry. *J.*
 839 *Atmospheric Ocean. Technol.* **31**, 181–196 (2014).
- 840 107. Golden Software, L. (2025).
- 841 108. Schlitzer, R. Ocean Data View. (2024).
- 842 109. De Ruggiero, P. *et al.* Modelling the marine circulation of the Campania coastal system
 843 (Tyrrhenian Sea) for the year 2016: Analysis of the dynamics. *J. Mar. Syst.* **210**, 103388 (2020).
- 844 110. Oddo, P. *et al.* A nested Atlantic-Mediterranean Sea general circulation model for operational
 845 forecasting. *Ocean Sci.* **5**, 461–473 (2009).
- 846 111. Madec, G. & Nemo team. *NEMO Ocean Engine*. (2008).
- 847 112. Pisano, A., Buongiorno Nardelli, B., Tronconi, C. & Santoleri, R. The new Mediterranean
 848 optimally interpolated pathfinder AVHRR SST Dataset (1982–2012). *Remote Sens. Environ.* **176**,
 849 107–116 (2016).
- 850 113. Andrews, S. FastQC: a quality control tool for high throughput sequence data. (2010).
- 851 114. Menzel, P., Ng, K. L. & Krogh, A. Fast and sensitive taxonomic classification for
 852 metagenomics with Kaiju. *Nat. Commun.* **7**, 11257 (2016).
- 853 115. Chen, S., Zhou, Y., Chen, Y. & Gu, J. fastp: an ultra-fast all-in-one FASTQ preprocessor.
 854 *Bioinformatics* **34**, i884–i890 (2018).
- 855 116. Hera, M. R. *et al.* Metagenomic functional profiling: to sketch or not to sketch?
 856 *Bioinformatics* **40**, ii165–ii173 (2024).
- 857 117. McMurdie, P. J. & Holmes, S. phyloseq: an R package for reproducible interactive analysis
 858 and graphics of microbiome census data. *PLoS One* **8**, e61217 (2013).
- 859 118. Oksanen, J. *et al.* vegan: Community Ecology Package. Ordination methods, diversity analysis
 860 and other functions for community and vegetation ecologists. Version 2.4-6. URL [https://CRAN.R-](https://CRAN.R-project.org/package=vegan)
 861 [project.org/package=vegan](https://CRAN.R-project.org/package=vegan). (2018).
- 862 119. Wickham, H. ggplot2. *Wiley Interdiscip. Rev. Comput. Stat.* **3**, 180–185 (2011).
- 863

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878

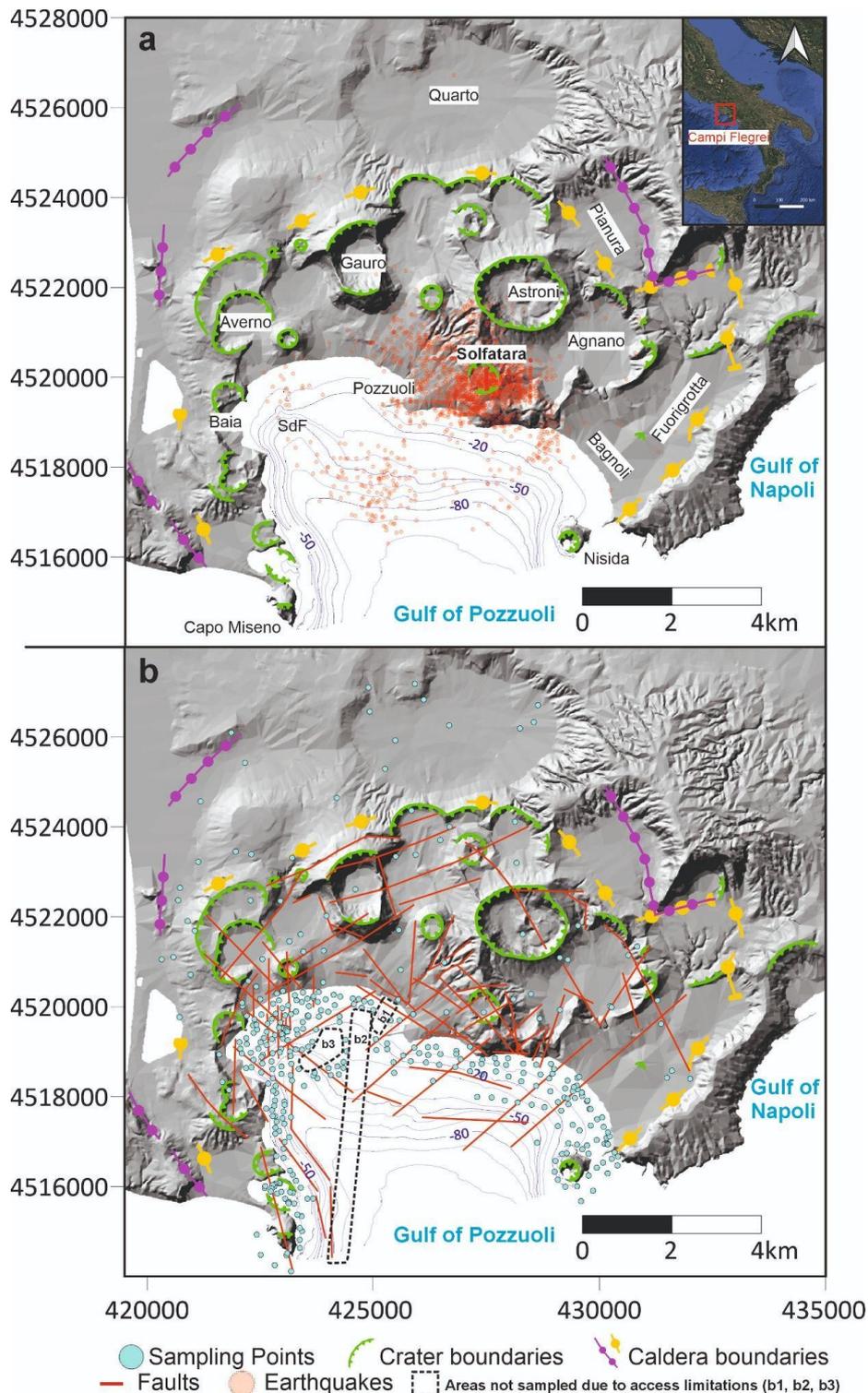
879 **Author contributions**

880 A.S.V., F.M., S.C.D., S.C., M.A.D.V., D.G., and A.A., drafted the original version of the manuscript.
881 A.A., D.G., S.C. and M.M conceived the study design. A.S.V., F.M., S.C.D and S.C. contributed to
882 geochemical and microbiological data collection and analysis. GS generated the ocean circulation
883 model. All coauthors contributed to editing the final version of the manuscript.

884 **Competing interests**

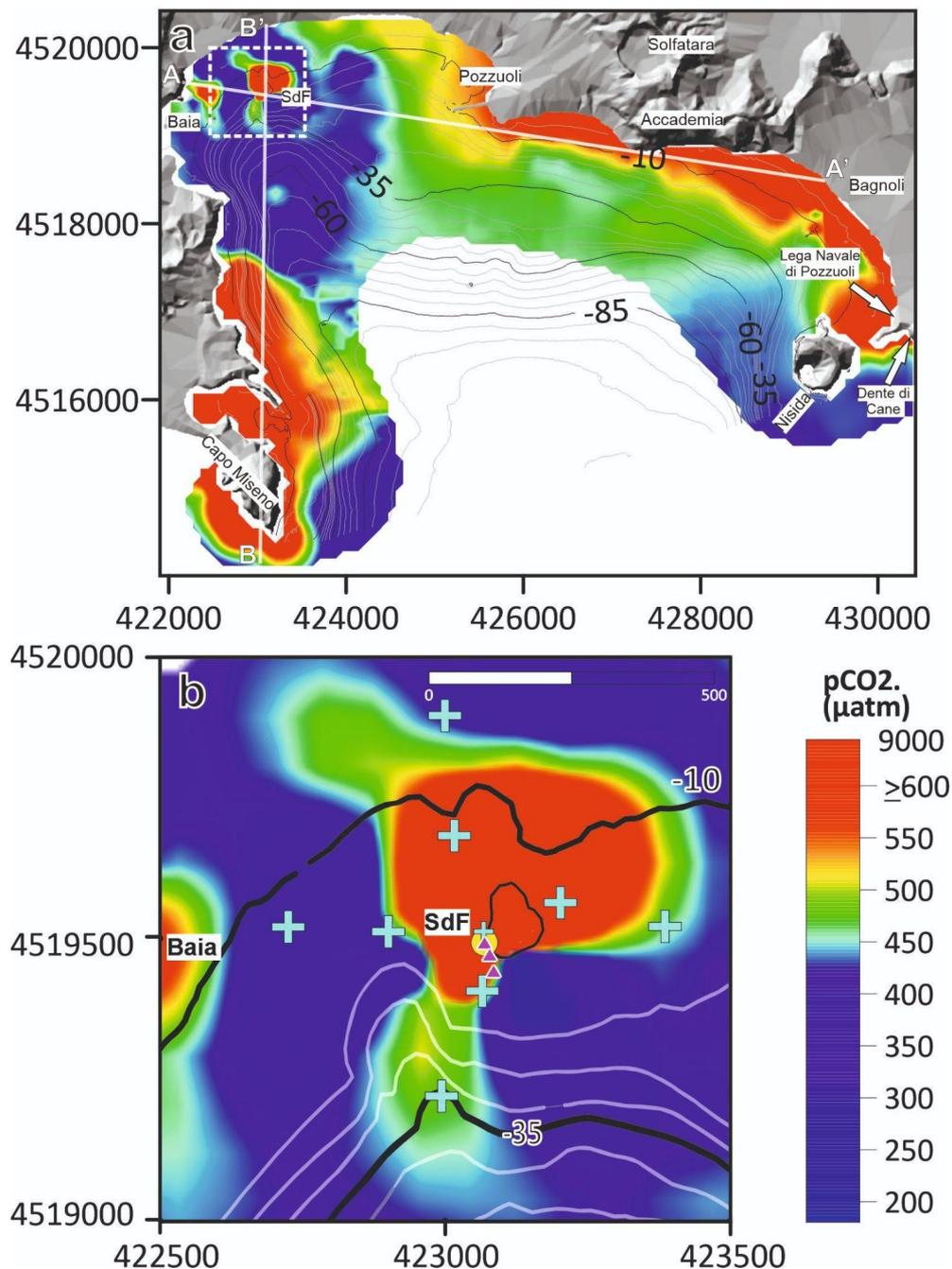
885 The authors declare no competing interests.

886 **Figures**



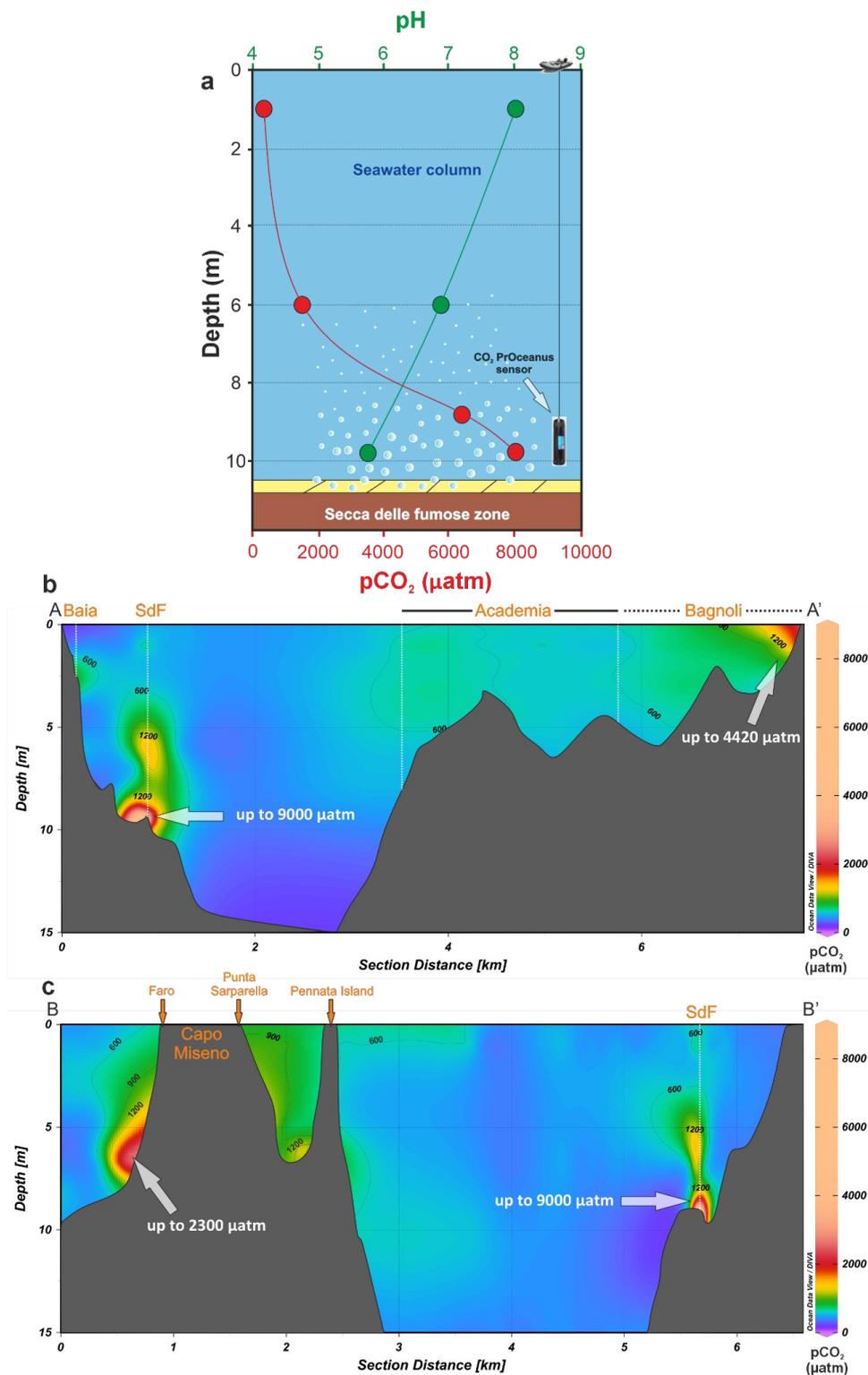
887

888 **Figure 1. Map of Campi Flegrei caldera (CFc).** These maps show the main localities, main geological
 889 structures^{36,67,73}, seismicity after 2022²⁷, bathymetry (given in meters) and sampling points. Areas not
 890 sampled due to access limitations are: b1) Pozzuoli harbour, b2) high-traffic navigation corridor and
 891 b3) the mytilid aquaculture facility. SdF: Secca delle Fumose. The image at the top right of panel a was
 892 taken from Google Earth Pro version 7.3.6.10441.



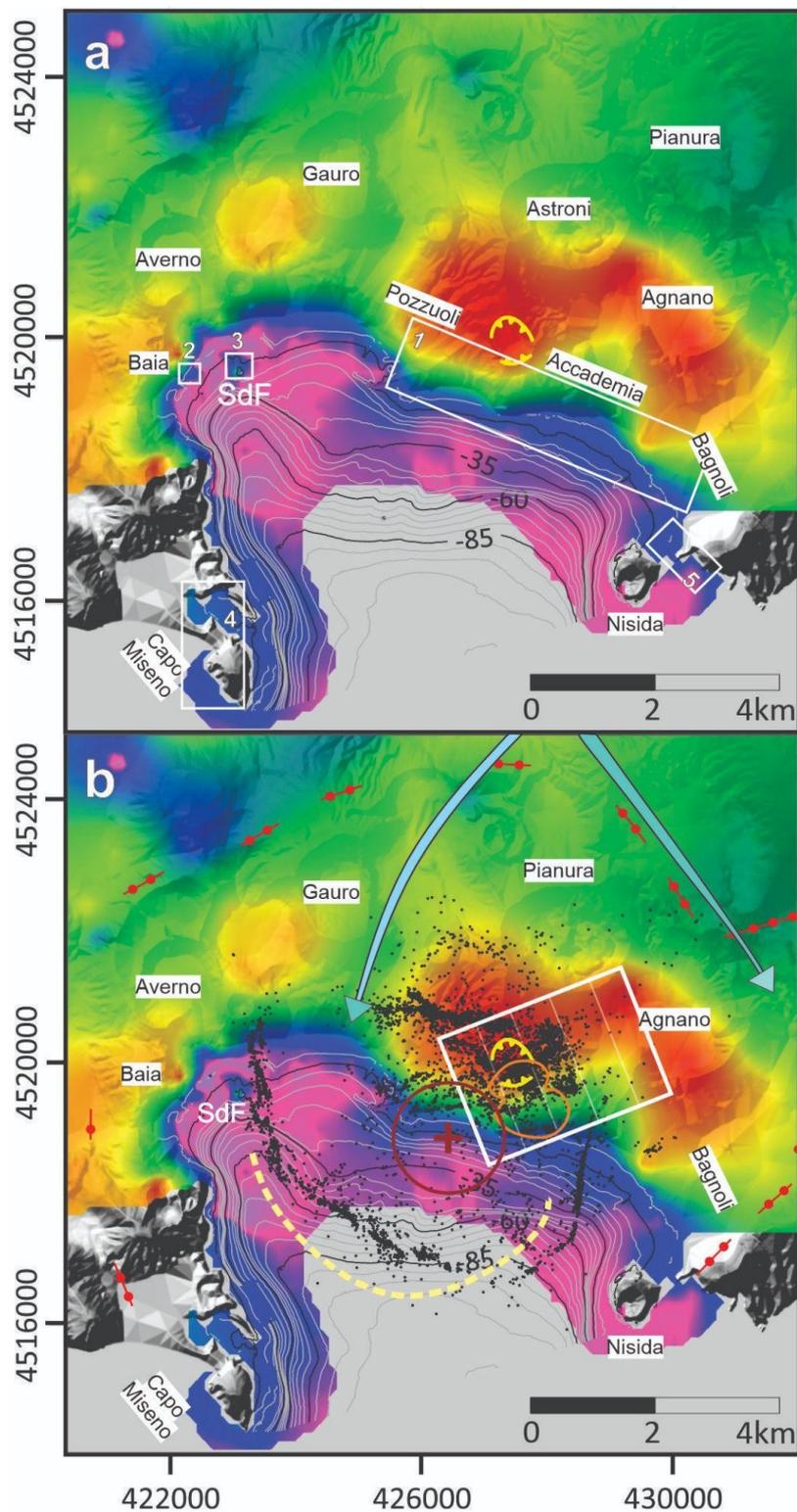
893

894 **Figure 2. Distribution of pCO₂ at the base of the seawater column along the Pozzuoli Bay.** a) The
 895 map was generated using the maximum pCO₂ value (pCO₂ max) measured at the deepest point of each
 896 station (see Supplementary Tables 1 and 2). Major anomalies were found in SdF (up to 9000 µatm),
 897 Accademia-Bagnoli-Nisida coast (up to 4420 µatm), Baia Castle beach (up to 850 µatm) and Capo
 898 Miseno (up to 2300 µatm). Black and gray solid lines represent bathymetry which is given in meters.
 899 Solid white lines represent the profiles A-A' and B-B' of Fig. 3. The white square represents the location
 900 of panel b, which is a closer view of the SdF anomaly. In this area, pCO₂ probing was coupled with the
 901 collection of biological samples, including biofilms (violet triangle), vent fluids (yellow circle), and
 902 seawater (blue cross). Samples from seawater column were collected along two transects, one N-S (sites
 903 T1, T2, T4, T5) and one W-E (sites T6, T7, T8, T9) centered on (site T3).



904

905 **Figure 3. Vertical variability of pCO₂ and pH in seawater above the SdF zone.** a) A profile showing
 906 the change of pCO₂ and pH directly above the SdF geyser using data presented in Supplementary Table
 907 2. b-c) Profiles showing the variability of pCO₂ in the seawater column and the distribution of the main
 908 hydrothermal hotspots along the Pozzuoli Gulf. Profiles AA' and BB' can be also observed in the map
 909 of Fig. 2. Data interpolation was achieved using DIVA Gridding available in ODV software (further
 910 details are presented in the “Methods” section).

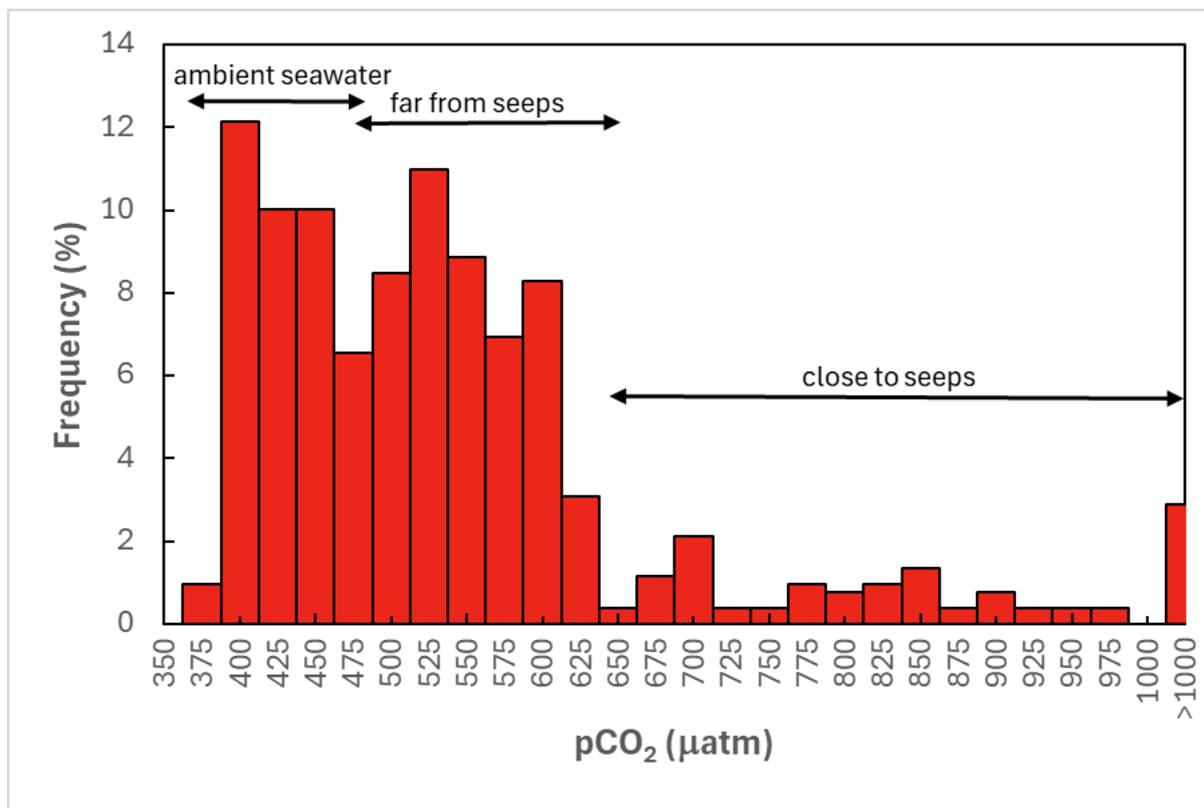


911

912 **Figure 4. Maps of CFC showing the pCO₂ distribution in seawater (Pozzuoli Bay) and**
 913 **groundwater. a) Regarding the Pozzuoli Bay, the map was generated using the maximum pCO₂ value**
 914 **(pCO₂ max) measured at the deepest point of each station. White squares indicate the main sectors (1**
 915 **to 5) where pCO₂ is anonymously high in seawater. b) Map of CFC comparing the pCO₂ distribution in**
 916 **water and the main morpho-structural elements observed in the caldera: The red lines (with red circles)**
 917 **represent the outer caldera rim bordering faults and the solid yellow lines the boundary of the Solfatara**

918 crater^{67,25}; the white square and the white lines indicate the orientation (N 249° ± 4°, 53° ± 1°) of the
 919 potential incipient volcanotectonic fault (PIF)⁷⁰. The dotted pale yellow line represents the base of the
 920 resurgent dome while the red circle (and cross in the center) shows the area with the maximum
 921 resurgence³⁹; the orange polygon represents the Accademia geodetic anomaly⁷¹; the black dots indicates
 922 the seismicity after 2022²⁷ while the blue long arrows show the aquifer main preferential drainage in
 923 the Campi Flegrei area³². Both maps were generated using the data of Supplementary Tables 1 and 2).

924



925

926 **Figure 5. Frequency distribution diagram of seawater pCO₂ values in the CFc offshore. Data are**
 927 **interpreted as a mixture of 3 distinct populations.**

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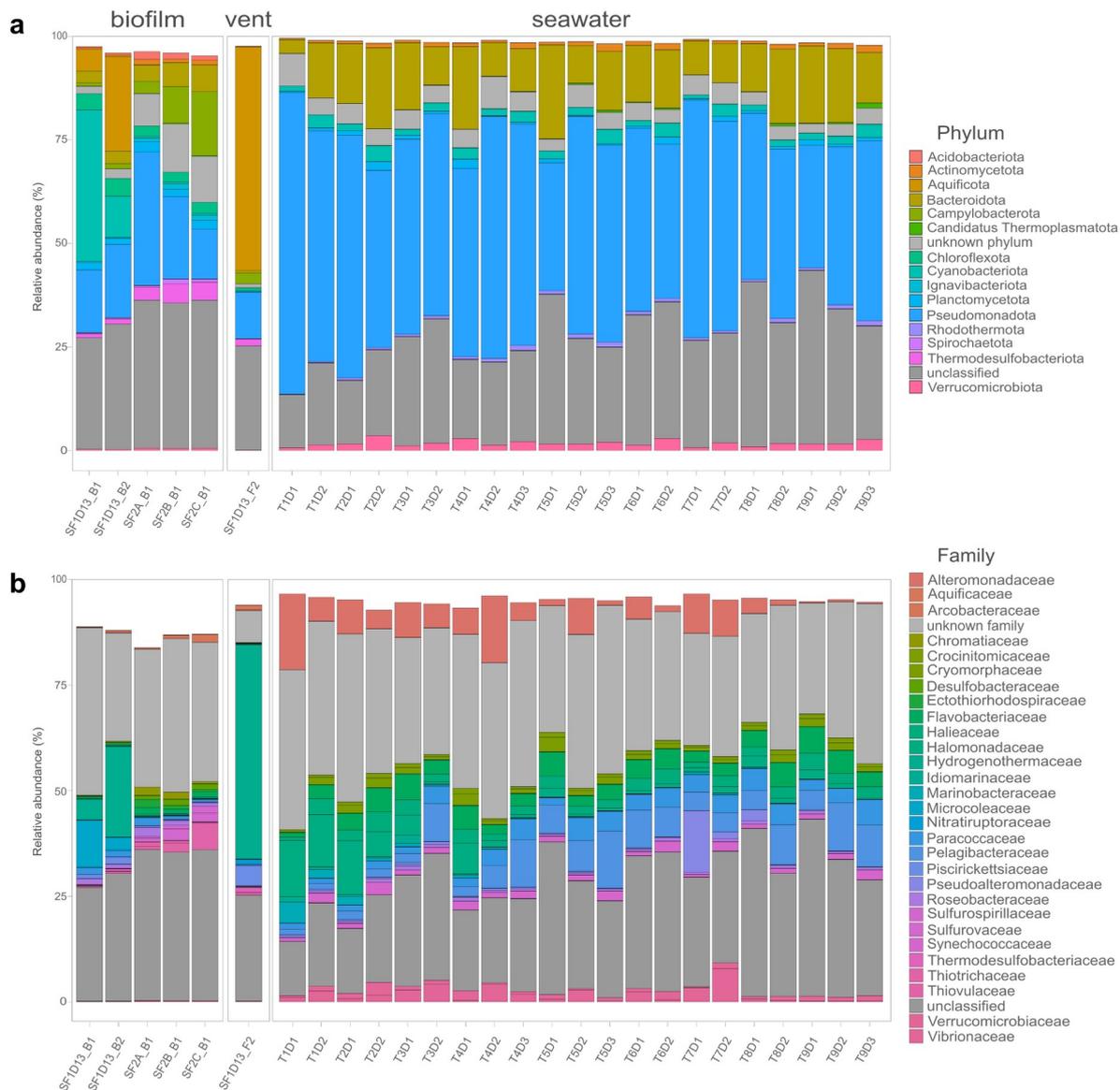
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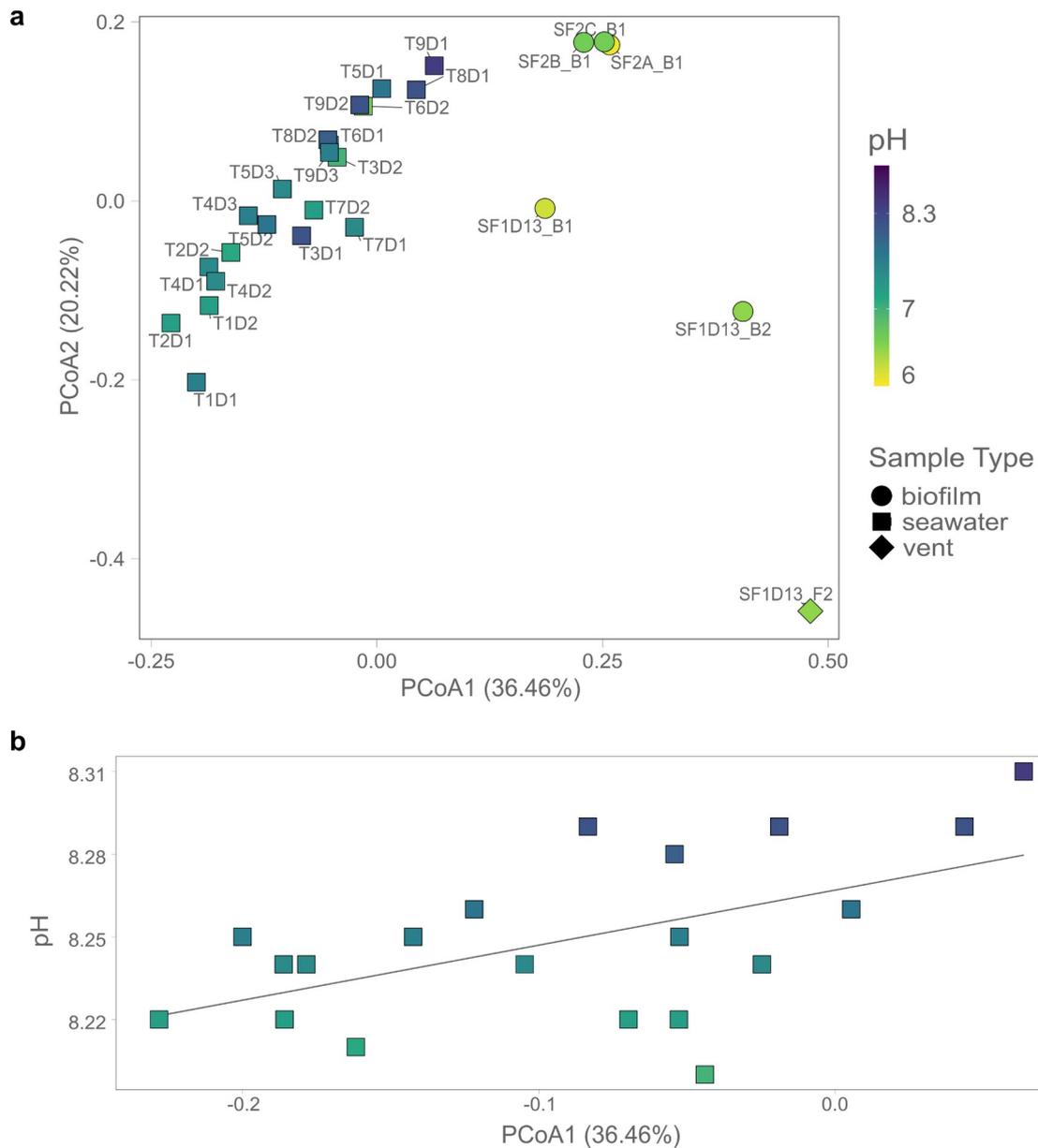


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937 **Figure 6. Microbial community composition across the SdF hydrothermal system.** (a) Relative
 938 taxonomic composition at the phylum level and (b) at the family level across biofilms, hydrothermal
 939 vent fluids, and seawater samples. Samples are grouped by environmental compartment, highlighting
 940 compositional contrasts between vent-associated communities and the surrounding seawater.

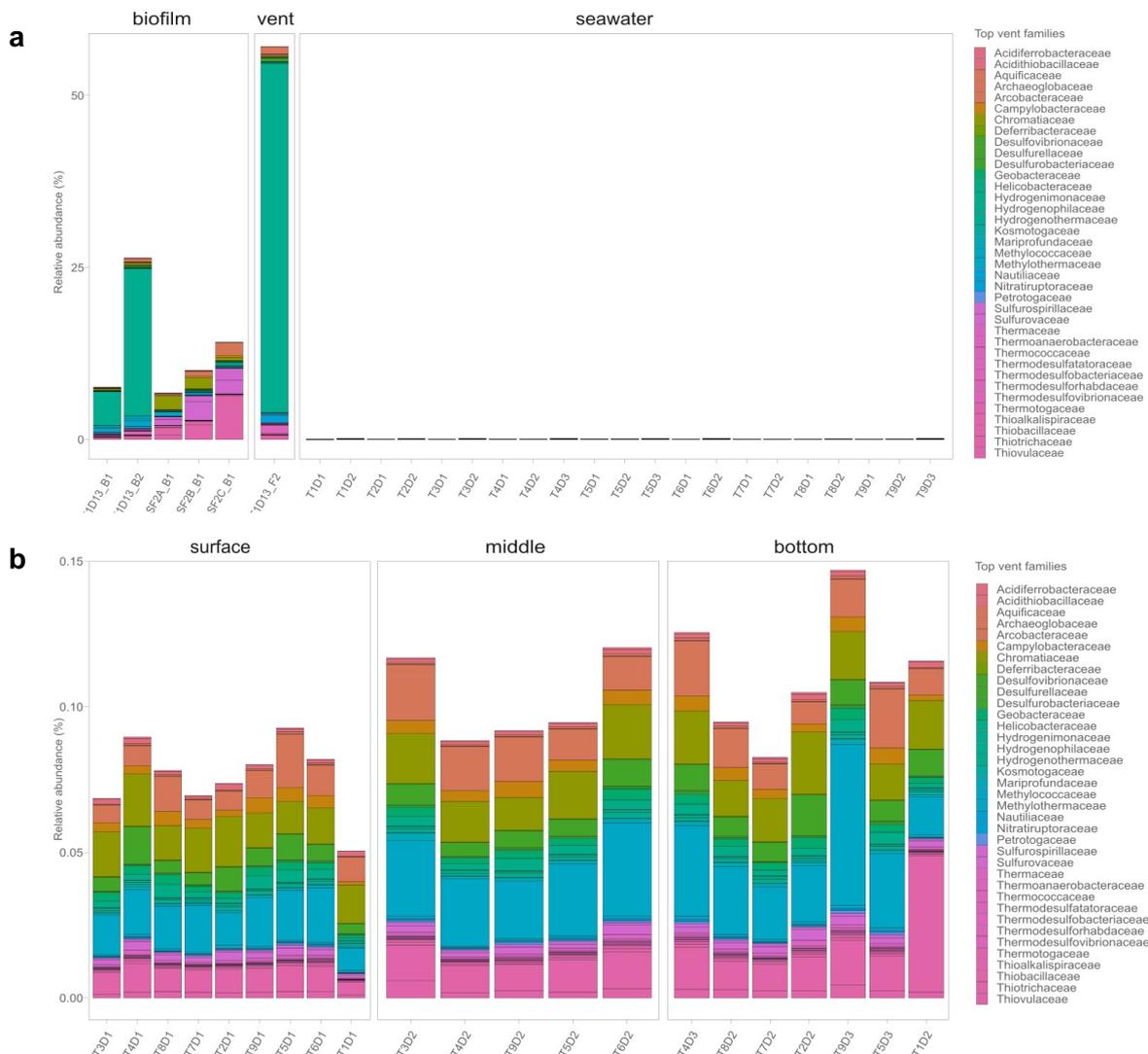
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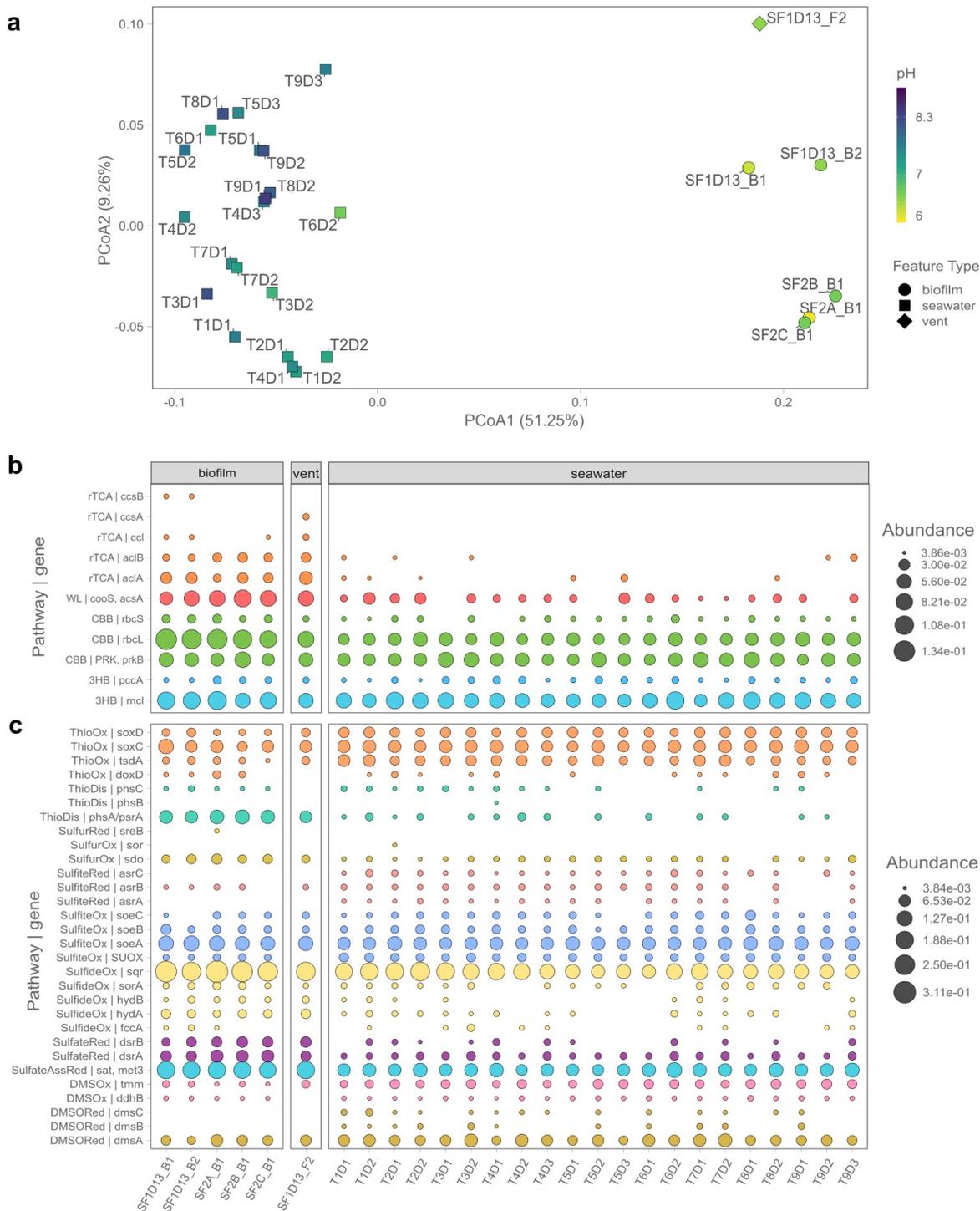


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944 **Figure 7. Beta diversity patterns of microbial communities in the SdF area.** (a) Principal
 945 Coordinates Analysis (PCoA) based on weighted Jaccard dissimilarity at the genus level. Symbols are
 946 shaped by sample type (biofilms, circles; seawater, squares; vent fluids, diamonds) and colored by pH.
 947 (b) Relationship between PCoA axis 1 and pH for seawater samples (Pearson correlation, $r^2 = 0.549$, p
 948 $= 0.012$; $n=20$).



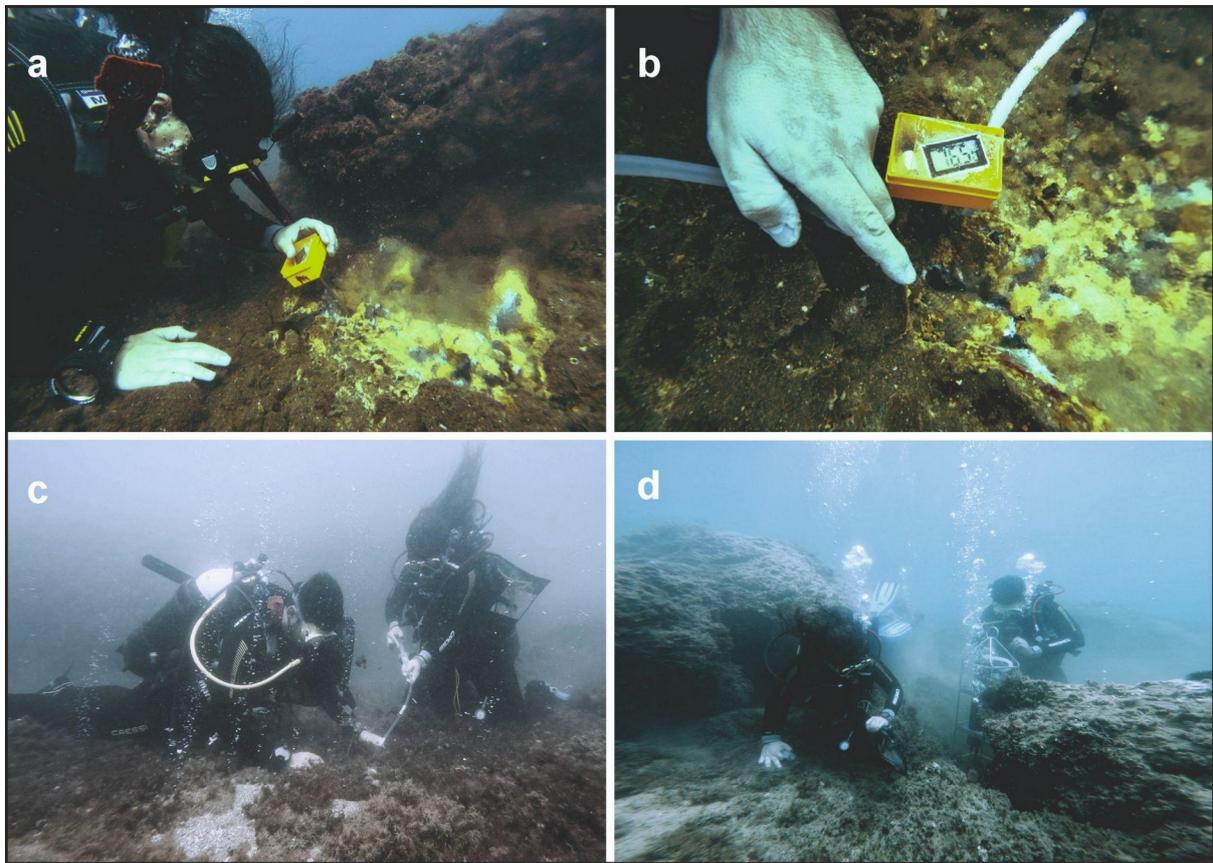
951 **Figure 8. Distribution of vent-associated microbial families in the SdF area.** (a) Relative abundance
 952 of selected vent-associated families across biofilm, hydrothermal vent fluid, and seawater samples. (b)
 953 Seawater samples shown by depth layer (surface, middle, bottom) and ordered by increasing distance
 954 from the SdF vent within each depth, highlighting dilution patterns away from the source.



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961 **Figure 9. Functional potential of microbial communities across the SdF hydrothermal system.** (a)
 962 Principal Coordinates Analysis (PCoA) based on weighted Jaccard dissimilarity of functional gene
 963 profiles. Symbols are shaped by sample type (biofilms, circles; seawater, squares; vent fluids,
 964 diamonds) and colored by pH. (b) Bubble plot showing the relative abundance of key genes involved
 965 in carbon fixation pathways (rTCA, reductive tricarboxylic acid cycle; WL, Wood–Ljungdahl pathway;
 966 CBB, Calvin–Benson–Bassham cycle; 3HB, 3-hydroxypropionate bicycle) across samples. (c) Bubble

967 plot showing the relative abundance of key genes involved in sulfur metabolism (thiosulfate oxidation
968 and disproportionation; sulfur, sulfite, sulfide, and sulfate oxidation and reduction; and dimethyl sulfide
969 oxidation and reduction). In panels B and C, bubble size is proportional to gene relative abundance.

970 **Supplementary Materials**971 **Supplementary Figures**

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973 **Supplementary Figure 1. Geochemical and microbiological sampling at Pozzuoli bay. SdF area (a,**
974 **b) and Capo Miseno (c, d) near El Faro (see Fig. 3c). Photos by Francesco Guerra used with**
975 **permission.**

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980 **Supplementary Figure 2. A close up view of the Secca delle Fumose main emission site**
981 **(SF1D13). The orange and green biofilm are visible.**

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988 **Supplementary Figure 3.** Area of lower temperature degassing near Secca delle Fumose.

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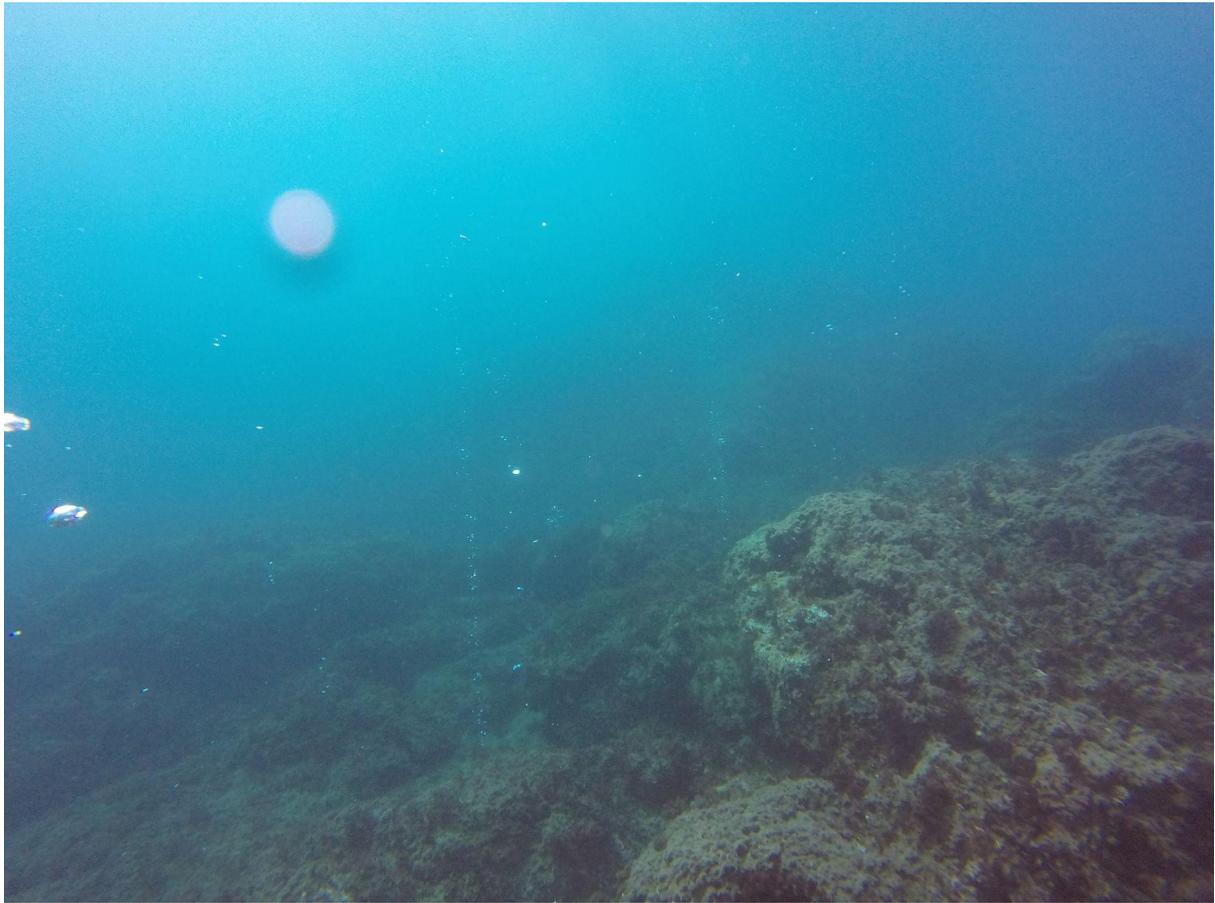
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993 **Supplementary Figure 4. Area of low temperature emission at Secca delle Fumose. Low**
994 **temperature white biofilms (SF2 samples) in the**

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999 **Supplementary Figure 5.** Presence of visible degassing in the area of Capo Miseno.

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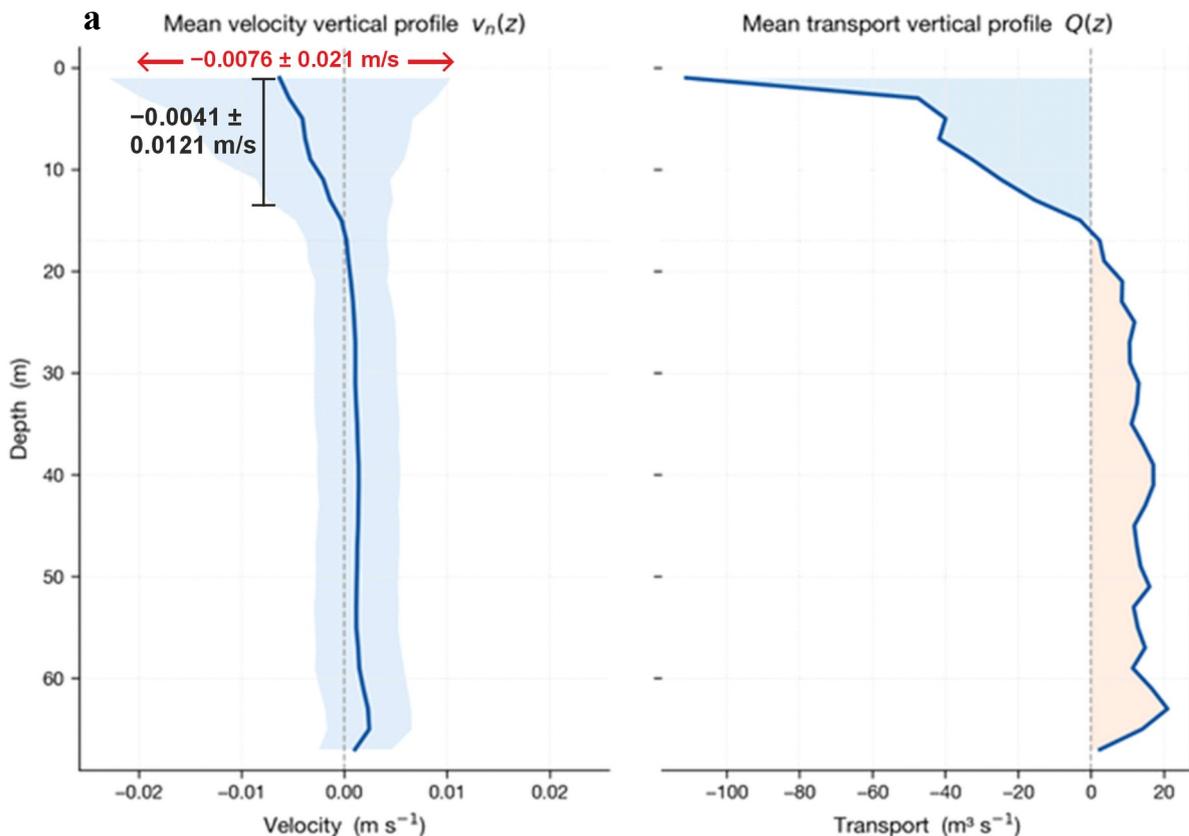
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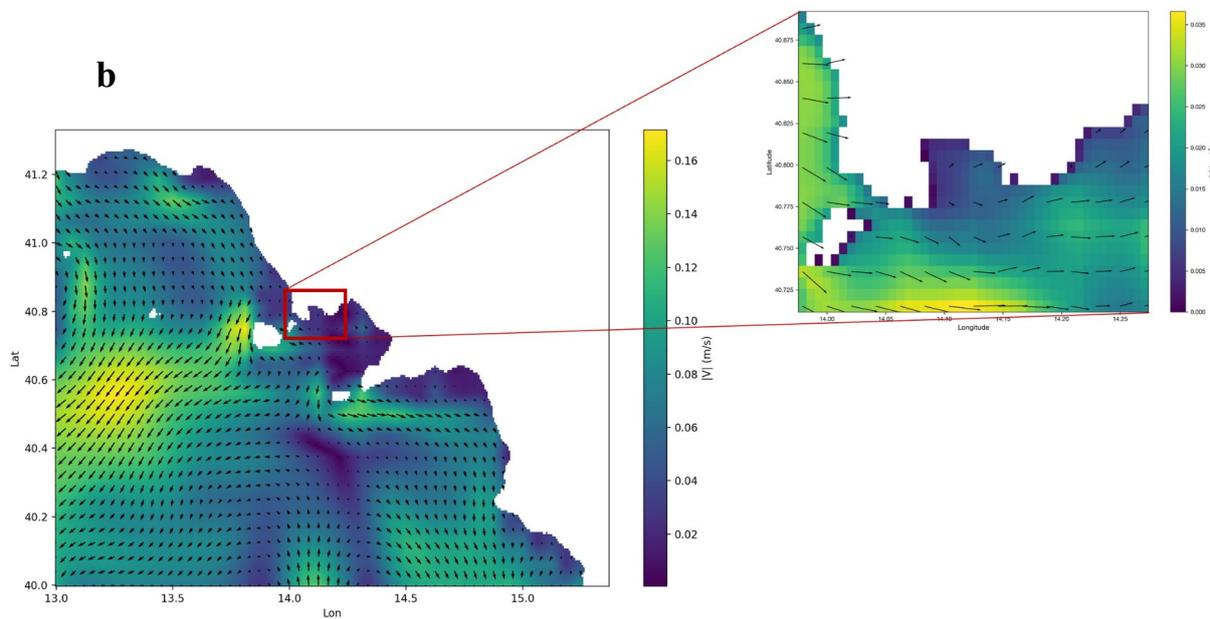
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Supplementary Figure 6: Vertical profile of annual mean velocity (in m/s) and transport (in m³/s) across the Baia–Bagnoli transect. a) Positive values indicate flow out of the gulf normal to the section, whereas negative values indicate flow into the gulf. Values were converted from sigma coordinates to physical depths and averaged along the section using area-weighted binning. Shading in the left panel

1035 represents the standard deviation of the annual mean velocity. Shadings in the right panel represents the
1036 net landward (blue) and seaward (red) transport. Statistics were computed after excluding the spin-up
1037 period. b) Surface velocity field showing the study area and a zoomed-in view of the Gulf of Pozzuoli.
1038 Colors represent current speed and arrows indicate flow direction.

Supplementary Table 1. pCO₂ values obtained from Seawater Vertical Profiles in the Pozzuoli Bay and from given in meters below sea level (b.s.l). Excess of carbon for each seawater point was calculated in reference to

Station	East X	North Y	Bottom Depth [m]	Depth [m]	pCO ₂ Max (μatm)	pCO ₂ (error +/-)
BG	423146	4518995.7	18	2	498	2.5
BG	423146	4518995.7	18	7	477	2.4
BG	423146	4518995.7	18	15	446	2.2
29	423089	4519440.4	15	2	513	2.6
29	423089	4519440.4	15	7	512	2.6
29	423089	4519440.4	15	13	518	2.6
SF1	423069	4519492.4	10	2	514	2.6
SF1	423069	4519492.4	10	7	1428	7.1
39	423095	4519503.3	11	2	522	2.6
45	422993	4519388.9	15	3	575	2.9
45	422993	4519388.9	15	8	529	2.6
46	422990	4519285.6	15	3	529	2.6
46	424158	4630303.2	15	7	523	2.6
46	425350	4741339.6	15	13	528	2.6
47	423006	4519514.2	11	2	534	2.7
47	423006	4519514.2	11	6	533	2.7
47	423006	4519514.2	11	10	532	2.7
48	423105	4519594.2	8	2	577	2.9
48	423105	4519594.2	8	7	532	2.7
54	422180	4519297.5	4	3	460	2.3
56	422494	4519203.1	10	1	423	2.1
56	422494	4519203.1	10	9	477	2.4
58	422626	4519144	14.3	6	425	2.1
58	422626	4519144	14.3	12	434	2.2
62	422569	4519301.1	10	5	422	2.1
62	422569	4519301.1	10	9	421	2.1
66	422327	4519579	3	3	849	4.2
70	422418	4519488.1	7	1	452	2.3
70	422418	4519488.1	7	6	549	2.7
72	422629	4519438.2	10	5	452	2.3
76	422904	4519527.9	12	1	438	2.2
86	422742	4519725.6	7	1	480	2.4
88	422934	4519645.9	11	1	470	2.4
88	422934	4519645.9	11	5	506	2.5
88	422934	4519645.9	11	10	487	2.4
90	423036	4519802.5	9.3	1	518	2.6
90	423036	4519802.5	9.3	8	518	2.6
92	422827	4519838	6	5	536	2.7
96	422772	4520062.8	2	1	422	2.1
112	422047	4519388.8	3	2	470	2.4
119	422024	4519207	5	3	476	2.4
121	422187	4519147.5	5.6	3	437	2.2

142	422047	4519062.4	4	2	423	2.1
144	422078	4518898.9	5	3	436	2.2
154	422173	4518631.4	2	1	435	2.2
155	422290	4518643.5	8	1	433	2.2
156	422539	4518545.4	25	1	427	2.1
162	422660	4518281	5	2	427	2.1
164	422800	4518144.1	25	1	428	2.1
179	422616	4517730.8	6	3	431	2.2
184	423289	4520138.5	2	1	443	2.2
185	423251	4520248.8	3.6	2	450	2.3
186	423406	4520099.5	5.18	3	442	2.2
187	423430	4520312.4	2	1	457	2.3
188	423527	4520174.9	3.6	2	475	2.4
189	423562	4520008	5.2	3	445	2.2
190	423783	4519777	10	1	439	2.2
190	423783	4519777	10	4	445	2.2
190	423783	4519777	10	9	448	2.2
191	423831	4519915.3	6.6	3	468	2.3
192	423757	4520104.8	2	1	478	2.4
193	423654	4520190.2	2	1	483	2.4
194	423665	4520226.7	3	1	480	2.4
195	423426	4519651.9	10	1	460	2.3
195	423426	4519651.9	10	4	453	2.3
195	423426	4519651.9	10	9	451	2.3
196	423145	4519027.6	43	1	454	2.3
196	423145	4519027.6	43	7	438	2.2
198	423055	4518877.5	-	1	453	2.3
198	423055	4518877.5	-	7	436	2.2
200	423123	4518712.5	54.2	1	449	2.2
200	423123	4518712.5	54.2	7	427	2.1
202	423144	4518589	48.3	1	452	2.3
202	423144	4518589	48.3	7	426	2.1
204	423268	4518252.5	64	1	486	2.4
204	423268	4518252.5	64	7	435	2.2
204	423268	4518252.5	64	15	442	2.2
206	423455	4518232.8	60.9	1	475	2.4
206	423455	4518232.8	60.9	7	434	2.2
208	423387	4518529.9	60	1	480	2.4
208	423387	4518529.9	60	7	449	2.2
210	423398	4518548.7	59.9	1	480	2.4
210	423398	4518548.7	59.9	7	451	2.3
212	423433	4519503.1	14	1	502	2.5
212	423433	4519503.1	14	6	469	2.3
212	423433	4519503.1	14	13	423	2.1
213	423091	4519252.4	20	1	490	2.5
213	423091	4519252.4	20	7	456	2.3
213	423091	4519252.4	20	15	432	2.2

214	423078	4519437.9	9	1	473	2.4
214	423078	4519437.9	9	4	466	2.3
214	423078	4519437.9	9	8	428	2.1
216	422964	4519385.8	15	7	491	2.5
216	422964	4519385.8	15	14	519	2.6
217	423292	4519484.6	12	1	488	2.4
217	423292	4519484.6	12	5	462	2.3
217	423292	4519484.6	12	11	443	2.2
218	423312	4519377.8	15.3	1	487	2.4
218	423312	4519377.8	15.3	7	464	2.3
218	423312	4519377.8	15.3	14	429	2.1
219	423413	4519305.7	-	1	489	2.4
220	423552	4519796.1	8	1	500	2.5
220	423552	4519796.1	8	6	468	2.3
221	423931	4520157.4	3	1	491	2.5
222	423777	4520321.1	2.7	1	527	2.6
223	424109	4520259.9	3	1	479	2.4
224	424132	4520029.9	6	3	486	2.4
225	423189	4518355.4	61	1	507	2.5
225	423189	4518355.4	61	7	460	2.3
227	424046	4518474.3	45	1	498	2.5
227	424046	4518474.3	45	7	445	2.2
229	424224	4518490.2	38	1	496	2.5
229	424224	4518490.2	38	7	444	2.2
231	424327	4518632.4	29	1	496	2.5
231	424327	4518632.4	29	7	451	2.3
233	424394	4519005.8	27	1	487	2.4
233	424394	4519005.8	27	7	457	2.3
233	424394	4519005.8	27	15	426	2.1
236	424239	4519777.9	16	1	512	2.6
236	424239	4519777.9	16	7	436	2.2
237	424238	4519656.9	18	1	509	2.5
237	424238	4519656.9	18	7	481	2.4
237	424238	4519656.9	18	15	439	2.2
238	424255	4520219.6	3	1	505	2.5
241	425892	4519083.9	6	1	695	3.5
241	425892	4519083.9	6	5	620	3.1
242	426195	4519096	5	3	644	3.2
243	426543	4519099.6	3.2	1	657	3.3
244	426823	4519011.9	5	3	576	2.9
245	427100	4518938.7	6	1	675	3.4
245	427100	4518938.7	6	5	585	2.9
246	427398	4518815.4	7.2	1	617	3.1
246	427398	4518815.4	7.2	6	582	2.9
247	427677	4518849.4	5	3	599	3.0
248	427965	4518848.2	5.2	3	591	3.0
249	428302	4518727.7	6	1	687	3.4

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253	428585	4518232.7	13	1	598	3.0
253	428585	4518232.7	13	6	577	2.9
253	428585	4518232.7	13	12	560	2.8
254	428258	4518333.8	15	1	550	2.8
254	428258	4518333.8	15	6	538	2.7
254	428258	4518333.8	15	12	522	2.6
255	427978	4518442.3	13	1	542	2.7
255	427978	4518442.3	13	6	539	2.7
255	427978	4518442.3	13	12	513	2.6
256	427698	4518556.6	11	1	539	2.7
256	427698	4518556.6	11	5	541	2.7
256	427698	4518556.6	11	9	539	2.7
257	427383	4518486.7	18	1	575	2.9
257	427383	4518486.7	18	7	536	2.7
257	427383	4518486.7	18	15	482	2.4
258	427125	4518560.9	19	1	606	3.0
258	427125	4518560.9	19	7	514	2.6
258	427125	4518560.9	19	15	498	2.5
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259	426799	4518654.3	18	15	471	2.4
260	425176	4518717.4	20	1	614	3.1
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260	425176	4518717.4	20	15	479	2.4
261	426151	4518755.2	18	1	590	3.0
261	426151	4518755.2	18	7	505	2.5
261	426151	4518755.2	18	15	521	2.6
262	423085	4519442.6	10	10	695	3.5
263	423069	4519455.9	10	9	1188	5.9
264	423078	4519487.9	10.8	8	818	4.1
265	423054	4519467.4	9.3	1	693	3.5
269	425812	4518783.7	18	1	553	2.8
269	425812	4518783.7	18	7	549	2.7
269	425812	4518783.7	18	15	526	2.6
270	425433	4518761.6	25	1	548	2.7
270	425433	4518761.6	25	7	574	2.9
270	425433	4518761.6	25	15	520	2.6
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271	428772	4517757.8	25	15	513	2.6
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272	428491	4517813.1	28	7	506	2.5

272	428491	4517813.1	28	15	489	2.4
273	428177	4517862.9	31	1	548	2.7
273	428177	4517862.9	31	7	533	2.7
273	428177	4517862.9	31	15	470	2.4
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274	427892	4517908.6	33	15	489	2.4
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275	427603	4517975.7	34	15	490	2.5
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276	427281	4517987.7	33	15	506	2.5
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278	425446	4518523.1	31	15	476	2.4
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280	429212	4517974.5	12	9	516	2.6
281	429399	4518113.7	4.7	1	589	2.9
281	429399	4518113.7	4.7	3	568	2.8
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283	429484	4518241.7	2.4	1	4408	22.0
284	429422	4518210	3.7	1	815	4.1
285	429229	4518176.4	6.4	1	621	3.1
285	429229	4518176.4	6.4	5	566	2.8
286	429283	4518351.3	3.5	1	831	4.2
287	429055	4518283.5	7.4	1	658	3.3
287	429055	4518283.5	7.4	5	574	2.9
288	428901	4518111.8	12	1	604	3.0
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291	428682	4518488	7.7	1	746	3.7
291	428682	4518488	7.7	5	701	3.5
292	428308	4518521.6	9.5	1	594	3.0
292	428308	4518521.6	9.5	8	489	2.4
293	427923	4518698.6	7.4	1	546	2.7
293	427923	4518698.6	7.4	5	504	2.5
294	427715	4518733.9	8.5	1	554	2.8
294	427715	4518733.9	8.5	6	512	2.6
295	427378	4518723.9	9.4	1	505	2.5

295	427378	4518723.9	9.4	7	500	2.5
296	427178	4518709.2	11	1	545	2.7
296	427178	4518709.2	11	8	519	2.6
297	426916	4518829.5	8.1	1	535	2.7
297	426916	4518829.5	8.1	6	510	2.5
298	426470	4518942.7	7.5	1	533	2.7
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300	426002	4518961.8	8.1	1	506	2.5
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301	425646	4518908.7	12	8	501	2.5
302	425294	4519051.1	18	1	508	2.5
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303	425291	4519173.2	13	8	516	2.6
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304	425040	4519273.4	17	8	496	2.5
305	425053	4519032.4	24	1	506	2.5
305	425053	4519032.4	24	8	499	2.5
306	425566	4519692.2	3.2	1	533	2.7
307	425370	4519689.8	3	1	541	2.7
308	424984	4519922.4	9	1	514	2.6
308	424984	4519922.4	9	6	497	2.5
309	424929	4520090.6	4.3	1	501	2.5
310	424789	4520115.3	2.8	1	516	2.6
311	424662	4520163.2	3	1	509	2.5
312	424531	4520180.1	2.9	1	516	2.6
313	424560	4520052.1	6.1	1	520	2.6
314	424729	4519970.5	8.7	1	519	2.6
314	424729	4519970.5	8.7	6	502	2.5
315	423066	4514466.6	8	1	590	2.9
315	423066	4514466.6	8	8	660	3.3
316	423084	4514510.7	7	6.7	2283	11.4
317	423361	4514627.3	16	1	610	3.0
317	423361	4514627.3	16	15	598	3.0
318	423358	4514809.2	9	1	589	2.9
318	423358	4514809.2	9	8	583	2.9
319	423410	4515002.1	8.5	1	582	2.9
319	423410	4515002.1	8.5	7.5	582	2.9
320	423361	4515071.6	6	1	549	2.7
320	423361	4515071.6	6	5	549	2.7
321	423399	4515334.8	10.6	1	574	2.9
321	423399	4515334.8	10.6	9	552	2.8
322	423306	4515367.4	8.7	1	553	2.8
322	423306	4515367.4	8.7	8	556	2.8

323	423236	4515615.4	8.5	1	577	2.9
323	423236	4515615.4	8.5	8	579	2.9
324	423529	4515722.2	13.5	1	592	3.0
324	423529	4515722.2	13.5	12	543	2.7
325	423209	4515858.3	6	1	690	3.5
325	423209	4515858.3	6	4.5	681	3.4
326	422923	4516060.4	5	1	761	3.8
326	422923	4516060.4	5	4.5	767	3.8
327	422910	4516071.1	5	1	796	4.0
327	422910	4516071.1	5	4.5	750	3.8
328	422836	4516066.3	5	1	789	3.9
328	422836	4516066.3	5	4	799	4.0
329	422805	4516043.6	7	1	829	4.1
329	422805	4516043.6	7	6	1676	8.4
330	422793	4515994.2	7.5	1	896	4.5
330	422793	4515994.2	7.5	6	842	4.2
331	422763	4516027.1	7	1	953	4.8
331	422763	4516027.1	7	6	825	4.1
332	422735	4515947.7	7	1	949	4.7
332	422735	4515947.7	7	5	894	4.5
333	422720	4515887.5	6.5	1	945	4.7
333	422720	4515887.5	6.5	5	1321	6.6
334	422587	4515961.2	4.5	1	958	4.8
335	422596	4516003.1	5	1	903	4.5
335	422596	4516003.1	5	4	796	4.0
336	422860	4515838	8.1	1	893	4.5
336	422860	4515838	8.1	6	844	4.2
337	422953	4515879.6	7.5	1	902	4.5
337	422953	4515879.6	7.5	5	676	3.4
338	423148	4515869.1	6.5	1	828	4.1
338	423148	4515869.1	6.5	4.5	681	3.4
339	423040	4515721.8	10.2	1	854	4.3
339	423040	4515721.8	10.2	7.5	675	3.4
340	423113	4515647.6	7	1	1055	5.3
340	423113	4515647.6	7	5	865	4.3
341	423112	4515621.7	6.6	1	1138	5.7
342	423656	4515725	20	1	588	2.9
342	423656	4515725	20	10	545	2.7
343	423382	4515969.8	12.5	1	589	2.9
343	423382	4515969.8	12.5	10	571	2.9
344	423277	4516121.9	12	1	608	3.0
344	423277	4516121.9	12	10	584	2.9
345	423205	4516255.9	11.4	1	601	3.0
345	423205	4516255.9	11.4	10	587	2.9
346	423143	4516350.9	10.5	1	590	2.9
346	423143	4516350.9	10.5	9.5	595	3.0
347	423065	4516506	11	1	606	3.0

347	423065	4516506	11	10	603	3.0
348	423001	4516633.3	10	1	597	3.0
348	423001	4516633.3	10	9	573	2.9
349	422999	4516833.1	32	1	588	2.9
349	422999	4516833.1	32	10	576	2.9
350	423027	4516986	41	1	582	2.9
350	423027	4516986	41	10	575	2.9
351	423125	4517092.7	48	1	584	2.9
351	423125	4517092.7	48	10	569	2.8
352	422902	4517228.2	42	1	583	2.9
352	422902	4517228.2	42	10	572	2.9
353	422820	4517170.3	16	1	582	2.9
353	422820	4517170.3	16	10	583	2.9
354	422823	4517303.5	20	1	583	2.9
354	422823	4517303.5	20	10	543	2.7
355	422849	4517440.8	23	1	581	2.9
355	422849	4517440.8	23	10	558	2.8
356	422817	4517561.1	18	1	584	2.9
356	422817	4517561.1	18	10	551	2.8
357	422763	4517561.6	14.5	1	581	2.9
357	422763	4517561.6	14.5	10	560	2.8
358	423017	4517391.4	47	1	589	2.9
358	423017	4517391.4	47	10	550	2.7
359	422980	4517512.8	44	1	580	2.9
359	422980	4517512.8	44	10	548	2.7
360	430383	4516463.5	6.5	1	447	2.2
360	430383	4516463.5	6.5	6	447	2.2
361	430260	4516612.3	8	1	470	2.4
361	430260	4516612.3	8	7	470	2.4
362	430067	4516526.4	14.5	1	547	2.7
362	430067	4516526.4	14.5	11	547	2.7
363	430320	4516755	2	1	1012	5.1
364	430398	4516663.2	3	1	571	2.9
365	429926	4516413.4	1	1	450	2.3
366	430166	4516287.9	12.5	1	432	2.2
367	429746	4516238.6	14	1	430	2.2
368	429844	4515900.2	22	1	430	2.2
369	429493	4516000.1	22	1	431	2.2
370	429580	4515667.3	28	1	431	2.2
371	429231	4515991.5	20	1	429	2.1
372	429067	4516147.4	30	1	434	2.2
373	428993	4516324.6	22	1	439	2.2
374	429014	4516566.4	13	1	440	2.2
375	429014	4516566.4	13	12	440	2.2
376	429156	4516716.1	15	1	444	2.2
377	429264	4516736.1	9	1	456	2.3
377	429264	4516736.1	9	8	456	2.3

378	428610	4516968.9	54	1	435	2.2
379	429297	4516955.6	20	1	480	2.4
380	429491	4516930.5	13	1	570	2.9
380	429491	4516930.5	13	10	570	2.9
381	429607	4516825	7	1	571	2.9
382	429742	4516724.9	8.5	1	613	3.1
383	429742	4516724.9	8.5	7.5	613	3.1
384	429713	4516909.5	7.5	1	590	3.0
385	429925	4517030.7	5	1	598	3.0
386	430101	4516945.8	3.5	1	645	3.2
387	430158	4516949.7	3.5	1	697	3.5
388	430173	4516932.9	3	1	762	3.8
389	430143	4517075.3	3	1	695	3.5
390	430210	4517111.3	2	1	814	4.1
391	430174	4517200.4	2.5	1	815	4.1
392	430079	4517321.2	4.5	1	556	2.8
393	430047	4517531.3	3.7	1	624	3.1
394	429935	4517491.3	5	1	523	2.6
395	429794	4517350.6	9	1	551	2.8
396	429548	4517243	12	1	532	2.7
397	429395	4517365.5	20	1	490	2.5
398	429067	4517339.7	23	1	449	2.2
399	429139	4517683.2	20	1	456	2.3
400	429301	4517824.9	10	1	475	2.4
401	429442	4517965.6	7	1	551	2.8
402	429604	4518012.9	3.5	1	766	3.8
403	429598	4518088.5	2.5	1	655	3.3
404	429713	4517889.8	3	1	1823	9.1
405	429722	4517910.8	2	1	1508	7.5
406	429767	4517721.6	7	1	525	2.6
407	429789	4517717	2.5	1	492	2.5
CF01	431479	4518414	-	-	36041	-
CF02	431588	4518570	-	-	22009	-
CF03	431270	4520135	-	-	19888	-
CF04	429495	4518411	-	-	82167	-
CF05	429540	4518436	-	-	126794	-
CF06	429513	4521799	-	-	5883	-
CF07	430627	4521138	-	-	7229	-
CF08	429288	4521042	-	-	781448	-
CF09	430820	4521335	-	-	3828	-
CF10	430220	4520007	-	-	798730	-
CF11	430254	4519984	-	-	836180	-
CF12	431350	4519613	-	-	73553	-
CF13	428473	4520279	-	-	217470	-
CF15	426673	4523998	-	-	1720	-
CF16	425538	4520154	-	-	31688	-
CF17	422990	4520347	-	-	5415	-

CF18	425591	4520818	-	-	35876	-
CF19	425923	4523376	-	-	22667	-
CF20	431993	4518392	-	-	18493	-
CF21	423442	4521220	-	-	85822	-
CF22	423951	4522018	-	-	92257	-
CF23	422092	4520008	-	-	37351	-
CF24	422085	4520080	-	-	49386	-
CF25	428290	4521610	-	-	19670	-
CF26	422224	4520568	-	-	52143	-
CF27	422399	4520443	-	-	24132	-
CF28	421576	4518931	-	-	23719	-
CF29	422854	4520288	-	-	61306	-
CF30	422934	4522349	-	-	41162	-
CF31	422270	4522360	-	-	5168	-
CF32	426470	4523703	-	-	6227	-
CF33	428156	4523197	-	-	7452	-
CF34	425248	4525288	-	-	1197	-
CF35	425473	4525921	-	-	42540	-
CF36	426102	4526828	-	-	29854	-
CF37	428230	4526188	-	-	13354	-
CF38	424907	4526562	-	-	17322	-
CF39	428628	4526704	-	-	25340	-
CF40	428535	4526327	-	-	9382	-
CF41	424880	4527098	-	-	993	-
CF42	421167	4524564	-	-	334	-
CF43	424180	4524642	-	-	1733	-
CF44	421116	4523218	-	-	15915	-
CF45	422291	4523389	-	-	474	-
CF46	421944	4523450	-	-	1870	-
CF47	422150	4525423	-	-	18022	-
CF48	424590	4522610	-	-	17898	-
CF49	422399	4520443	-	-	14686	-
CF50	425867	4519829	-	-	56938	-
CF51	425867	4519829	-	-	163192	-
CF52	423130	4521450	-	-	106121	-
CF53	423888	4521240	-	-	39039	-
CF54	422187	4519790	-	-	345541	-
CF55	428120	4524113	-	-	18264	-
CF56	425904	4527179	-	-	5226	-
CF57	423378	4521605	-	-	48652	-
CF58	423191	4523308	-	-	2150	-
CF59	425551	4521700	-	-	8582	-
CF60	420878	4521902	-	-	1172	-
CF61	420947	4520709	-	-	14545	-
CF62	421038	4521685	-	-	7532	-
CF63A	421845	4526093	-	-	4373	-
CF64	420701	4519764	-	-	132190	-

CF65	420713	4522640	-	-	10242	-
CF66	420308	4521050	-	-	8339	-
CF67	420566	4521099	-	-	14801	-
CF68	426603	4519873	-	-	180219	-
CF69	426662	4526255	-	-	80039	-
CF70	426557	4520976	-	-	482281	-
CF72	428087	4520160	-	-	144677	-
CF74	422533	4521056	-	-	7881	-
CF75	427427	4519870	-	-	867961	-
CF76	429114	4519924	-	-	25334	-
CF77	428670	4519730	-	-	19702	-
CF78	421795	4518144	-	-	597035	-
CF79	423876	4520537	-	-	1290	-
CF80	429110	4521088	-	-	392103	-
CF81	421650	4518322	-	-	5090	-
CF82	427122	4519107	-	-	5749	-
CF83	422988	4520348	-	-	317	-
CF84	425832	4524362	-	-	21429	-
CF85	425482	4523338	-	-	12145	-
CF86	420671	4522364	-	-	31521	-

groundwaters after Caliro et al. (2025b). Reported errors for pCO₂ are 1σ uncertainties. The depth is the atmospheric equilibrium pCO₂ value of 420 μm.

Log₁₀ (pCO₂)	Excess carbon (mol/kg)	Excess carbon (μmol/kg)	Type	Reference
2.7	4.2E-05	41.7	Seawater	This work
2.7	3.1E-05	31.3	Seawater	This work
2.6	1.5E-05	14.9	Seawater	This work
2.7	4.9E-05	48.8	Seawater	This work
2.7	4.8E-05	48.3	Seawater	This work
2.7	5.1E-05	51.1	Seawater	This work
2.7	4.9E-05	49.3	Seawater	This work
3.2	2.6E-04	257.7	Seawater	This work
2.7	5.3E-05	52.9	Seawater	This work
2.8	7.5E-05	75.5	Seawater	This work
2.7	5.6E-05	56.1	Seawater	This work
2.7	5.6E-05	56.1	Seawater	This work
2.7	5.3E-05	53.4	Seawater	This work
2.7	5.6E-05	55.6	Seawater	This work
2.7	5.8E-05	58.3	Seawater	This work
2.7	5.8E-05	57.8	Seawater	This work
2.7	5.7E-05	57.4	Seawater	This work
2.8	7.6E-05	76.3	Seawater	This work
2.7	5.7E-05	57.4	Seawater	This work
2.7	2.2E-05	22.5	Seawater	This work
2.6	1.8E-06	1.8	Seawater	This work
2.7	3.1E-05	31.3	Seawater	This work
2.6	3.0E-06	3.0	Seawater	This work
2.6	8.2E-06	8.2	Seawater	This work
2.6	1.2E-06	1.2	Seawater	This work
2.6	5.9E-07	0.6	Seawater	This work
2.9	1.6E-04	160.1	Seawater	This work
2.7	1.8E-05	18.2	Seawater	This work
2.7	6.5E-05	64.8	Seawater	This work
2.7	1.8E-05	18.2	Seawater	This work
2.6	1.0E-05	10.4	Seawater	This work
2.7	3.3E-05	32.8	Seawater	This work
2.7	2.8E-05	27.7	Seawater	This work
2.7	4.6E-05	45.5	Seawater	This work
2.7	3.6E-05	36.3	Seawater	This work
2.7	5.1E-05	51.1	Seawater	This work
2.7	5.1E-05	51.1	Seawater	This work
2.7	5.9E-05	59.2	Seawater	This work
2.6	1.2E-06	1.2	Seawater	This work
2.7	2.8E-05	27.7	Seawater	This work
2.7	3.1E-05	30.8	Seawater	This work
2.6	9.9E-06	9.9	Seawater	This work

2.6	1.8E-06	1.8	Seawater	This work
2.6	9.3E-06	9.3	Seawater	This work
2.6	8.7E-06	8.7	Seawater	This work
2.6	7.6E-06	7.6	Seawater	This work
2.6	4.1E-06	4.1	Seawater	This work
2.6	4.1E-06	4.1	Seawater	This work
2.6	4.7E-06	4.7	Seawater	This work
2.6	6.5E-06	6.5	Seawater	This work
2.6	1.3E-05	13.2	Seawater	This work
2.7	1.7E-05	17.1	Seawater	This work
2.6	1.3E-05	12.7	Seawater	This work
2.7	2.1E-05	20.9	Seawater	This work
2.7	3.0E-05	30.3	Seawater	This work
2.6	1.4E-05	14.4	Seawater	This work
2.6	1.1E-05	11.0	Seawater	This work
2.6	1.4E-05	14.4	Seawater	This work
2.7	1.6E-05	16.0	Seawater	This work
2.7	2.7E-05	26.7	Seawater	This work
2.7	3.2E-05	31.8	Seawater	This work
2.7	3.4E-05	34.3	Seawater	This work
2.7	3.3E-05	32.8	Seawater	This work
2.7	2.2E-05	22.5	Seawater	This work
2.7	1.9E-05	18.7	Seawater	This work
2.7	1.8E-05	17.6	Seawater	This work
2.7	1.9E-05	19.3	Seawater	This work
2.6	1.0E-05	10.4	Seawater	This work
2.7	1.9E-05	18.7	Seawater	This work
2.6	9.3E-06	9.3	Seawater	This work
2.7	1.7E-05	16.6	Seawater	This work
2.6	4.1E-06	4.1	Seawater	This work
2.7	1.8E-05	18.2	Seawater	This work
2.6	3.5E-06	3.5	Seawater	This work
2.7	3.6E-05	35.8	Seawater	This work
2.6	8.7E-06	8.7	Seawater	This work
2.6	1.3E-05	12.7	Seawater	This work
2.7	3.0E-05	30.3	Seawater	This work
2.6	8.2E-06	8.2	Seawater	This work
2.7	3.3E-05	32.8	Seawater	This work
2.7	1.7E-05	16.6	Seawater	This work
2.7	3.3E-05	32.8	Seawater	This work
2.7	1.8E-05	17.6	Seawater	This work
2.7	4.4E-05	43.6	Seawater	This work
2.7	2.7E-05	27.2	Seawater	This work
2.6	1.8E-06	1.8	Seawater	This work
2.7	3.8E-05	37.8	Seawater	This work
2.7	2.0E-05	20.4	Seawater	This work
2.6	7.0E-06	7.0	Seawater	This work

2.7	2.9E-05	29.3	Seawater	This work
2.7	2.6E-05	25.7	Seawater	This work
2.6	4.7E-06	4.7	Seawater	This work
2.7	3.8E-05	38.3	Seawater	This work
2.7	5.2E-05	51.6	Seawater	This work
2.7	3.7E-05	36.8	Seawater	This work
2.7	2.4E-05	23.6	Seawater	This work
2.6	1.3E-05	13.2	Seawater	This work
2.7	3.6E-05	36.3	Seawater	This work
2.7	2.5E-05	24.6	Seawater	This work
2.6	5.3E-06	5.3	Seawater	This work
2.7	3.7E-05	37.3	Seawater	This work
2.7	4.3E-05	42.7	Seawater	This work
2.7	2.7E-05	26.7	Seawater	This work
2.7	3.8E-05	38.3	Seawater	This work
2.7	5.5E-05	55.2	Seawater	This work
2.7	3.2E-05	32.3	Seawater	This work
2.7	3.6E-05	35.8	Seawater	This work
2.7	4.6E-05	46.0	Seawater	This work
2.7	2.2E-05	22.5	Seawater	This work
2.7	4.2E-05	41.7	Seawater	This work
2.6	1.4E-05	14.4	Seawater	This work
2.7	4.1E-05	40.7	Seawater	This work
2.6	1.4E-05	13.8	Seawater	This work
2.7	4.1E-05	40.7	Seawater	This work
2.7	1.8E-05	17.6	Seawater	This work
2.7	3.6E-05	36.3	Seawater	This work
2.7	2.1E-05	20.9	Seawater	This work
2.6	3.5E-06	3.5	Seawater	This work
2.7	4.8E-05	48.3	Seawater	This work
2.6	9.3E-06	9.3	Seawater	This work
2.7	4.7E-05	46.9	Seawater	This work
2.7	3.3E-05	33.3	Seawater	This work
2.6	1.1E-05	11.0	Seawater	This work
2.7	4.5E-05	45.0	Seawater	This work
2.8	1.2E-04	117.9	Seawater	This work
2.8	9.3E-05	92.7	Seawater	This work
2.8	1.0E-04	101.2	Seawater	This work
2.8	1.1E-04	105.6	Seawater	This work
2.8	7.6E-05	75.9	Seawater	This work
2.8	1.1E-04	111.6	Seawater	This work
2.8	7.9E-05	79.5	Seawater	This work
2.8	9.2E-05	91.6	Seawater	This work
2.8	7.8E-05	78.3	Seawater	This work
2.8	8.5E-05	84.9	Seawater	This work
2.8	8.2E-05	81.8	Seawater	This work
2.8	1.2E-04	115.4	Seawater	This work

2.8	8.6E-05	86.0	Seawater	This work
2.9	1.2E-04	122.0	Seawater	This work
2.9	1.3E-04	129.1	Seawater	This work
2.8	1.1E-04	109.9	Seawater	This work
2.7	6.9E-05	69.4	Seawater	This work
2.8	8.4E-05	84.5	Seawater	This work
2.8	7.6E-05	76.3	Seawater	This work
2.7	6.9E-05	69.4	Seawater	This work
2.7	6.5E-05	65.2	Seawater	This work
2.7	6.0E-05	60.0	Seawater	This work
2.7	5.3E-05	52.9	Seawater	This work
2.7	6.2E-05	61.8	Seawater	This work
2.7	6.0E-05	60.5	Seawater	This work
2.7	4.9E-05	48.8	Seawater	This work
2.7	6.0E-05	60.5	Seawater	This work
2.7	6.1E-05	61.3	Seawater	This work
2.7	6.0E-05	60.5	Seawater	This work
2.8	7.5E-05	75.5	Seawater	This work
2.7	5.9E-05	59.2	Seawater	This work
2.7	3.4E-05	33.8	Seawater	This work
2.8	8.7E-05	87.5	Seawater	This work
2.7	4.9E-05	49.3	Seawater	This work
2.7	4.2E-05	41.7	Seawater	This work
2.8	8.9E-05	88.6	Seawater	This work
2.7	4.7E-05	47.4	Seawater	This work
2.7	2.8E-05	28.3	Seawater	This work
2.8	9.0E-05	90.5	Seawater	This work
2.7	4.3E-05	42.7	Seawater	This work
2.7	3.2E-05	32.3	Seawater	This work
2.8	8.1E-05	81.4	Seawater	This work
2.7	4.5E-05	45.0	Seawater	This work
2.7	5.2E-05	52.5	Seawater	This work
2.8	1.2E-04	117.9	Seawater	This work
3.1	2.2E-04	225.0	Seawater	This work
2.9	1.5E-04	152.5	Seawater	This work
2.8	1.2E-04	117.3	Seawater	This work
2.7	6.6E-05	66.5	Seawater	This work
2.7	6.5E-05	64.8	Seawater	This work
2.7	5.5E-05	54.7	Seawater	This work
2.7	6.4E-05	64.3	Seawater	This work
2.8	7.5E-05	75.1	Seawater	This work
2.7	5.2E-05	52.0	Seawater	This work
2.8	7.9E-05	79.5	Seawater	This work
2.7	6.7E-05	67.3	Seawater	This work
2.7	4.9E-05	48.8	Seawater	This work
2.7	6.6E-05	66.5	Seawater	This work
2.7	4.6E-05	45.5	Seawater	This work

2.7	3.7E-05	37.3	Seawater	This work
2.7	6.4E-05	64.3	Seawater	This work
2.7	5.8E-05	57.8	Seawater	This work
2.7	2.8E-05	27.7	Seawater	This work
2.7	6.5E-05	64.8	Seawater	This work
2.7	5.2E-05	52.0	Seawater	This work
2.7	3.7E-05	37.3	Seawater	This work
2.7	6.9E-05	68.6	Seawater	This work
2.7	5.8E-05	58.3	Seawater	This work
2.7	3.8E-05	37.8	Seawater	This work
2.7	6.7E-05	67.3	Seawater	This work
2.7	6.0E-05	59.6	Seawater	This work
2.7	4.6E-05	45.5	Seawater	This work
2.8	7.3E-05	72.7	Seawater	This work
2.7	6.6E-05	66.0	Seawater	This work
2.7	3.2E-05	31.8	Seawater	This work
2.7	6.4E-05	63.9	Seawater	This work
2.7	5.7E-05	57.4	Seawater	This work
2.7	3.1E-05	30.8	Seawater	This work
2.7	6.9E-05	69.3	Seawater	This work
2.7	4.3E-05	42.9	Seawater	This work
2.7	3.7E-05	36.9	Seawater	This work
2.7	5.8E-05	58.0	Seawater	This work
2.7	5.0E-05	50.2	Seawater	This work
2.8	8.1E-05	80.9	Seawater	This work
2.8	7.2E-05	72.5	Seawater	This work
3.4	3.5E-04	345.8	Seawater	This work
3.6	4.4E-04	444.1	Seawater	This work
2.9	1.5E-04	151.6	Seawater	This work
2.8	9.3E-05	93.0	Seawater	This work
2.8	7.2E-05	71.9	Seawater	This work
2.9	1.6E-04	155.6	Seawater	This work
2.8	1.1E-04	105.8	Seawater	This work
2.8	7.5E-05	75.3	Seawater	This work
2.8	8.7E-05	86.9	Seawater	This work
2.7	5.5E-05	55.3	Seawater	This work
2.9	1.5E-04	151.3	Seawater	This work
2.9	1.7E-04	167.9	Seawater	This work
2.9	1.3E-04	133.2	Seawater	This work
2.8	1.2E-04	119.9	Seawater	This work
2.8	8.3E-05	83.0	Seawater	This work
2.7	3.7E-05	37.3	Seawater	This work
2.7	6.3E-05	63.4	Seawater	This work
2.7	4.5E-05	44.5	Seawater	This work
2.7	6.7E-05	66.9	Seawater	This work
2.7	4.8E-05	48.3	Seawater	This work
2.7	4.5E-05	44.9	Seawater	This work

2.7	4.3E-05	42.5	Seawater	This work
2.7	6.3E-05	62.9	Seawater	This work
2.7	5.2E-05	51.7	Seawater	This work
2.7	5.9E-05	58.8	Seawater	This work
2.7	4.7E-05	47.3	Seawater	This work
2.7	5.8E-05	57.9	Seawater	This work
2.7	4.7E-05	47.3	Seawater	This work
2.7	4.8E-05	48.4	Seawater	This work
2.7	4.4E-05	43.7	Seawater	This work
2.7	4.5E-05	45.4	Seawater	This work
2.7	4.5E-05	45.1	Seawater	This work
2.7	4.5E-05	45.1	Seawater	This work
2.7	4.3E-05	43.2	Seawater	This work
2.7	4.7E-05	46.6	Seawater	This work
2.7	4.6E-05	46.2	Seawater	This work
2.7	4.8E-05	47.9	Seawater	This work
2.7	5.0E-05	50.2	Seawater	This work
2.7	5.1E-05	51.2	Seawater	This work
2.7	4.1E-05	40.6	Seawater	This work
2.7	4.5E-05	45.4	Seawater	This work
2.7	4.2E-05	42.4	Seawater	This work
2.7	5.8E-05	58.0	Seawater	This work
2.7	6.1E-05	61.4	Seawater	This work
2.7	4.9E-05	49.5	Seawater	This work
2.7	4.1E-05	41.1	Seawater	This work
2.7	4.3E-05	43.4	Seawater	This work
2.7	5.0E-05	50.1	Seawater	This work
2.7	4.7E-05	46.9	Seawater	This work
2.7	5.0E-05	50.1	Seawater	This work
2.7	5.2E-05	52.0	Seawater	This work
2.7	5.2E-05	51.6	Seawater	This work
2.7	4.3E-05	43.4	Seawater	This work
2.8	8.1E-05	81.3	Seawater	This work
2.8	1.1E-04	106.6	Seawater	This work
3.4	3.4E-04	335.2	Seawater	This work
2.8	8.9E-05	88.9	Seawater	This work
2.8	8.5E-05	84.6	Seawater	This work
2.8	8.1E-05	81.2	Seawater	This work
2.8	7.9E-05	78.8	Seawater	This work
2.8	7.8E-05	78.1	Seawater	This work
2.8	7.8E-05	78.2	Seawater	This work
2.7	6.5E-05	64.7	Seawater	This work
2.7	6.5E-05	64.9	Seawater	This work
2.8	7.5E-05	74.9	Seawater	This work
2.7	6.6E-05	66.0	Seawater	This work
2.7	6.6E-05	66.4	Seawater	This work
2.7	6.8E-05	67.8	Seawater	This work

2.8	7.6E-05	76.3	Seawater	This work
2.8	7.7E-05	77.2	Seawater	This work
2.8	8.2E-05	82.1	Seawater	This work
2.7	6.2E-05	62.0	Seawater	This work
2.8	1.2E-04	116.4	Seawater	This work
2.8	1.1E-04	113.4	Seawater	This work
2.9	1.4E-04	137.5	Seawater	This work
2.9	1.4E-04	139.1	Seawater	This work
2.9	1.5E-04	146.9	Seawater	This work
2.9	1.3E-04	134.3	Seawater	This work
2.9	1.5E-04	145.0	Seawater	This work
2.9	1.5E-04	147.5	Seawater	This work
2.9	1.6E-04	155.2	Seawater	This work
3.2	2.8E-04	284.9	Seawater	This work
3.0	1.7E-04	170.9	Seawater	This work
2.9	1.6E-04	158.4	Seawater	This work
3.0	1.8E-04	183.2	Seawater	This work
2.9	1.5E-04	154.3	Seawater	This work
3.0	1.8E-04	182.3	Seawater	This work
3.0	1.7E-04	170.6	Seawater	This work
3.0	1.8E-04	181.4	Seawater	This work
3.1	2.4E-04	244.0	Seawater	This work
3.0	1.8E-04	184.3	Seawater	This work
3.0	1.7E-04	172.5	Seawater	This work
2.9	1.5E-04	146.8	Seawater	This work
3.0	1.7E-04	170.3	Seawater	This work
2.9	1.6E-04	158.8	Seawater	This work
3.0	1.7E-04	172.3	Seawater	This work
2.8	1.1E-04	111.9	Seawater	This work
2.9	1.5E-04	154.9	Seawater	This work
2.8	1.1E-04	113.4	Seawater	This work
2.9	1.6E-04	161.3	Seawater	This work
2.8	1.1E-04	111.6	Seawater	This work
3.0	2.0E-04	202.8	Seawater	This work
2.9	1.6E-04	163.9	Seawater	This work
3.1	2.2E-04	217.1	Seawater	This work
2.8	8.1E-05	80.5	Seawater	This work
2.7	6.3E-05	62.9	Seawater	This work
2.8	8.1E-05	81.0	Seawater	This work
2.8	7.4E-05	74.0	Seawater	This work
2.8	8.8E-05	88.3	Seawater	This work
2.8	7.9E-05	79.2	Seawater	This work
2.8	8.5E-05	85.5	Seawater	This work
2.8	8.0E-05	80.1	Seawater	This work
2.8	8.1E-05	81.3	Seawater	This work
2.8	8.3E-05	83.2	Seawater	This work
2.8	8.7E-05	87.4	Seawater	This work

2.8	8.6E-05	86.5	Seawater	This work
2.8	8.4E-05	83.9	Seawater	This work
2.8	7.5E-05	74.6	Seawater	This work
2.8	8.1E-05	80.6	Seawater	This work
2.8	7.6E-05	76.0	Seawater	This work
2.8	7.8E-05	78.4	Seawater	This work
2.8	7.5E-05	75.4	Seawater	This work
2.8	7.9E-05	79.0	Seawater	This work
2.8	7.3E-05	73.2	Seawater	This work
2.8	7.9E-05	78.6	Seawater	This work
2.8	7.4E-05	74.2	Seawater	This work
2.8	7.8E-05	78.2	Seawater	This work
2.8	7.9E-05	78.7	Seawater	This work
2.8	7.9E-05	78.8	Seawater	This work
2.7	6.2E-05	62.2	Seawater	This work
2.8	7.8E-05	78.0	Seawater	This work
2.7	6.9E-05	68.5	Seawater	This work
2.8	7.9E-05	79.2	Seawater	This work
2.7	6.6E-05	65.6	Seawater	This work
2.8	7.8E-05	78.1	Seawater	This work
2.7	7.0E-05	69.6	Seawater	This work
2.8	8.1E-05	81.2	Seawater	This work
2.7	6.5E-05	65.1	Seawater	This work
2.8	7.8E-05	77.5	Seawater	This work
2.7	6.4E-05	64.3	Seawater	This work
2.7	1.5E-05	15.5	Seawater	This work
2.7	1.5E-05	15.5	Seawater	This work
2.7	2.8E-05	27.7	Seawater	This work
2.7	2.8E-05	27.7	Seawater	This work
2.7	6.4E-05	63.9	Seawater	This work
2.7	6.4E-05	63.9	Seawater	This work
3.0	1.9E-04	194.9	Seawater	This work
2.8	7.4E-05	73.9	Seawater	This work
2.7	1.7E-05	17.1	Seawater	This work
2.6	7.0E-06	7.0	Seawater	This work
2.6	5.9E-06	5.9	Seawater	This work
2.6	5.9E-06	5.9	Seawater	This work
2.6	6.4E-06	6.4	Seawater	This work
2.6	6.4E-06	6.4	Seawater	This work
2.6	5.3E-06	5.3	Seawater	This work
2.6	8.2E-06	8.2	Seawater	This work
2.6	1.1E-05	11.0	Seawater	This work
2.6	1.2E-05	11.6	Seawater	This work
2.6	1.2E-05	11.6	Seawater	This work
2.6	1.4E-05	13.8	Seawater	This work
2.7	2.0E-05	20.4	Seawater	This work
2.7	2.0E-05	20.4	Seawater	This work

2.6	8.7E-06	8.7	Seawater	This work
2.7	3.3E-05	32.8	Seawater	This work
2.8	7.3E-05	73.5	Seawater	This work
2.8	7.3E-05	73.5	Seawater	This work
2.8	7.4E-05	73.9	Seawater	This work
2.8	9.0E-05	90.1	Seawater	This work
2.8	9.0E-05	90.1	Seawater	This work
2.8	8.1E-05	81.4	Seawater	This work
2.8	8.4E-05	84.5	Seawater	This work
2.8	1.0E-04	101.5	Seawater	This work
2.8	1.2E-04	118.6	Seawater	This work
2.9	1.4E-04	137.7	Seawater	This work
2.8	1.2E-04	117.9	Seawater	This work
2.9	1.5E-04	151.4	Seawater	This work
2.9	1.5E-04	151.7	Seawater	This work
2.7	6.8E-05	67.7	Seawater	This work
2.8	9.4E-05	94.1	Seawater	This work
2.7	5.3E-05	53.4	Seawater	This work
2.7	6.6E-05	65.6	Seawater	This work
2.7	5.7E-05	57.4	Seawater	This work
2.7	3.8E-05	37.8	Seawater	This work
2.7	1.7E-05	16.6	Seawater	This work
2.7	2.0E-05	20.4	Seawater	This work
2.7	3.0E-05	30.3	Seawater	This work
2.7	6.6E-05	65.6	Seawater	This work
2.9	1.4E-04	138.8	Seawater	This work
2.8	1.0E-04	104.9	Seawater	This work
3.3	3.0E-04	298.8	Seawater	This work
3.2	2.7E-04	267.1	Seawater	This work
2.7	5.4E-05	54.3	Seawater	This work
2.7	3.9E-05	38.8	Seawater	This work
4.6	-	-	Groundwater	Caliro et al. (2025b)
4.3	-	-	Groundwater	Caliro et al. (2025b)
4.3	-	-	Groundwater	Caliro et al. (2025b)
4.9	-	-	Groundwater	Caliro et al. (2025b)
5.1	-	-	Groundwater	Caliro et al. (2025b)
3.8	-	-	Groundwater	Caliro et al. (2025b)
3.9	-	-	Groundwater	Caliro et al. (2025b)
5.9	-	-	Groundwater	Caliro et al. (2025b)
3.6	-	-	Groundwater	Caliro et al. (2025b)
5.9	-	-	Groundwater	Caliro et al. (2025b)
5.9	-	-	Groundwater	Caliro et al. (2025b)
4.9	-	-	Groundwater	Caliro et al. (2025b)
5.3	-	-	Groundwater	Caliro et al. (2025b)
3.2	-	-	Groundwater	Caliro et al. (2025b)
4.5	-	-	Groundwater	Caliro et al. (2025b)
3.7	-	-	Groundwater	Caliro et al. (2025b)

4.6	-	-	Groundwater Caliro et al. (2025b)
4.4	-	-	Groundwater Caliro et al. (2025b)
4.3	-	-	Groundwater Caliro et al. (2025b)
4.9	-	-	Groundwater Caliro et al. (2025b)
5.0	-	-	Groundwater Caliro et al. (2025b)
4.6	-	-	Groundwater Caliro et al. (2025b)
4.7	-	-	Groundwater Caliro et al. (2025b)
4.3	-	-	Groundwater Caliro et al. (2025b)
4.7	-	-	Groundwater Caliro et al. (2025b)
4.4	-	-	Groundwater Caliro et al. (2025b)
4.4	-	-	Groundwater Caliro et al. (2025b)
4.8	-	-	Groundwater Caliro et al. (2025b)
4.6	-	-	Groundwater Caliro et al. (2025b)
3.7	-	-	Groundwater Caliro et al. (2025b)
3.8	-	-	Groundwater Caliro et al. (2025b)
3.9	-	-	Groundwater Caliro et al. (2025b)
3.1	-	-	Groundwater Caliro et al. (2025b)
4.6	-	-	Groundwater Caliro et al. (2025b)
4.5	-	-	Groundwater Caliro et al. (2025b)
4.1	-	-	Groundwater Caliro et al. (2025b)
4.2	-	-	Groundwater Caliro et al. (2025b)
4.4	-	-	Groundwater Caliro et al. (2025b)
4.0	-	-	Groundwater Caliro et al. (2025b)
3.0	-	-	Groundwater Caliro et al. (2025b)
2.5	-	-	Groundwater Caliro et al. (2025b)
3.2	-	-	Groundwater Caliro et al. (2025b)
4.2	-	-	Groundwater Caliro et al. (2025b)
2.7	-	-	Groundwater Caliro et al. (2025b)
3.3	-	-	Groundwater Caliro et al. (2025b)
4.3	-	-	Groundwater Caliro et al. (2025b)
4.3	-	-	Groundwater Caliro et al. (2025b)
4.2	-	-	Groundwater Caliro et al. (2025b)
4.8	-	-	Groundwater Caliro et al. (2025b)
5.2	-	-	Groundwater Caliro et al. (2025b)
5.0	-	-	Groundwater Caliro et al. (2025b)
4.6	-	-	Groundwater Caliro et al. (2025b)
5.5	-	-	Groundwater Caliro et al. (2025b)
4.3	-	-	Groundwater Caliro et al. (2025b)
3.7	-	-	Groundwater Caliro et al. (2025b)
4.7	-	-	Groundwater Caliro et al. (2025b)
3.3	-	-	Groundwater Caliro et al. (2025b)
3.9	-	-	Groundwater Caliro et al. (2025b)
3.1	-	-	Groundwater Caliro et al. (2025b)
4.2	-	-	Groundwater Caliro et al. (2025b)
3.9	-	-	Groundwater Caliro et al. (2025b)
3.6	-	-	Groundwater Caliro et al. (2025b)
5.1	-	-	Groundwater Caliro et al. (2025b)

4.0	-	-	Groundwater Caliro et al. (2025b)
3.9	-	-	Groundwater Caliro et al. (2025b)
4.2	-	-	Groundwater Caliro et al. (2025b)
5.3	-	-	Groundwater Caliro et al. (2025b)
4.9	-	-	Groundwater Caliro et al. (2025b)
5.7	-	-	Groundwater Caliro et al. (2025b)
5.2	-	-	Groundwater Caliro et al. (2025b)
3.9	-	-	Groundwater Caliro et al. (2025b)
5.9	-	-	Groundwater Caliro et al. (2025b)
4.4	-	-	Groundwater Caliro et al. (2025b)
4.3	-	-	Groundwater Caliro et al. (2025b)
5.8	-	-	Groundwater Caliro et al. (2025b)
3.1	-	-	Groundwater Caliro et al. (2025b)
5.6	-	-	Groundwater Caliro et al. (2025b)
3.7	-	-	Groundwater Caliro et al. (2025b)
3.8	-	-	Groundwater Caliro et al. (2025b)
2.5	-	-	Groundwater Caliro et al. (2025b)
4.3	-	-	Groundwater Caliro et al. (2025b)
4.1	-	-	Groundwater Caliro et al. (2025b)
4.5	-	-	Groundwater Caliro et al. (2025b)

Supplementary Table 2. pCO₂ and pH values obtained from a seawater vertical profile performed immediately above the Fumose vent (SdF). Reported errors for pCO₂ are 1σ uncertainties. The depth is given in meters below sea level (b.s.l). Excess carbon for each seawater point was calculated in reference to the atmospheric equilibrium pCO₂ value of 420 μm.

Station	East X	North Y	Bottom Depth [m]	Depth [m]	pCO₂ Max (μatm)	pCO₂ (error +/-)	Excess carbon (mol/kg)	Excess carbon (μmol/kg)
SF2	423065	4519508	10	2	514	2.6	4.9E-05	49.3
SF2	423065	4519508	10	6	1560	7.8	2.7E-04	272.9
SF2	423065	4519508	10	10	6822	34.1	5.3E-04	532.2
SF2	423065	4519508	10	9.3	8028	40.1	5.7E-04	571.8

Secca delle
ess of

pH

- 8.10
- 6.90
- 6.05
- 5.90