

Spectral signatures in satellite soil moisture reveal irrigation patterns across the contiguous United States

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12 **Key Points:**

- 13 • Irrigation modifies the spectral characteristics of soil moisture.
- 14 • Spectral differences between SMOS observations and irrigation-off Noah–MP sim-
- 15 ulations detect irrigated areas.
- 16 • A SMOS-derived irrigation map improves the representation of irrigation patterns
- 17 in Noah–MP simulations across the contiguous United States.

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Abstract

Irrigation profoundly alters the terrestrial water cycle, yet its spatial distribution and temporal variability remain poorly constrained. Here, we introduce a new approach to detect irrigation in space based on spectral differences between modelled and satellite-observed soil moisture time series. Using wavelet decomposition, we isolate irrigation-induced variability at sub-annual scales by comparing Noah–MP simulations with and without irrigation. We then analyse differences between Soil Moisture and Ocean Salinity (SMOS) soil moisture retrievals and model simulations without irrigation to identify irrigated areas. We apply this method over the Contiguous United States (CONUS) for the period 2010–2023. Results show that irrigation enhances the power of seasonal and sub-seasonal spectral components of soil moisture, enabling identification of irrigated and non-irrigated areas when comparing observed and modelled soil moisture. Comparison with LGRIP30, an irrigation dataset for the United States, shows moderate-to-good agreement in the spatial patterns of irrigation. Major irrigated regions across CONUS largely coincide with those identified in LGRIP30, while discrepancies increase in areas with weaker irrigation signals or dense vegetation. The resulting SMOS-derived irrigation map suggests spatial patterns of irrigated areas that differ from those represented in static reference maps such as LGRIP30. When used to constrain Noah–MP simulations, SMOS-derived irrigation map highlights alternative irrigation distributions that better reflect actual irrigation spatial patterns and water use. Our method provides a scalable, all-weather pathway to detect where irrigation occurs, offering new observational constraints for representing irrigation in land and Earth-system models.

Plain Language Summary

Irrigation is essential for global food production, but we still do not know how much water is used for irrigation and how it varies over time and across regions. This limits our ability to understand the water cycle and to plan for sustainable water management. Satellite missions such as ESA’s Soil Moisture and Ocean Salinity (SMOS) provide regular global measurements of soil moisture, which can reveal where irrigation is happening.

In this study, we analyzed SMOS soil moisture data together with computer model simulations to look at the “spectral fingerprint” of irrigation – the way it changes soil moisture patterns over time. Using a mathematical technique called wavelet analysis,

50 we isolated the portion of the signal most influenced by irrigation and trained a machine
51 learning algorithm to identify irrigated areas.

52 Our results show that this approach can detect irrigation across the contiguous United
53 States with good accuracy. This method can be applied globally, offering new opportu-
54 nities to monitor irrigation and improve water management.

55 **1 Introduction**

56 Irrigation is one of the most significant human interventions in the terrestrial water
57 cycle, yet its distribution and extent remain highly uncertain, limiting our ability to
58 close the water balance and to represent land-atmosphere interactions in models (Mc-
59 Dermid et al., 2023). Reliable global information on where and when irrigation occurs
60 is challenging to obtain, and most land surface models rely on maps derived from sub-
61 national irrigation inventories that fail to capture its dynamic and heterogeneous nature
62 (Neumann et al., 2011).

63 Satellite observations provide a unique opportunity to detect and quantify irriga-
64 tion, using, for example, retrievals of soil moisture (Brocca et al., 2018), vegetation ac-
65 tivity (Chiesi et al., 2022), evapotranspiration (Brombacher et al., 2022), and land sur-
66 face temperature (Purnamasari et al., 2025), with near-global coverage and frequent re-
67 visit times (Massari et al., 2021). Specifically, some previous studies have attempted to
68 detect and quantify irrigation by comparing satellite soil moisture retrievals with model
69 outputs that exclude irrigation and analysing differences in their variability (Lawston
70 et al., 2017; Zaussinger et al., 2019). However, despite promising results in certain re-
71 gions, irrigation detection remains challenging due to uncertainties in the spatial foot-
72 print of satellite soil moisture retrievals and the presence of several confounding factors,
73 including topography, heterogeneous vegetation cover, agricultural practices such as tillage,
74 satellite signal disturbances such as radio frequency interference, and the often smaller
75 spatial scale of irrigation relative to the satellite footprint.

76 Here we present a new method that exploits the spectral signature of irrigation in
77 satellite soil moisture time series. Using wavelet decomposition, we isolate irrigation-driven
78 variability at sub-annual scales and derive spectral features from the differences between
79 simulations with and without irrigation. These features are used to train a random for-
80 est classifier on synthetic soil moisture time series generated with the Noah-MP land sur-

81 face model (Niu et al., 2011). The trained classifier is then applied to wavelet-transformed
82 soil moisture observations from SMOS (Kerr et al., 2010) to detect irrigated areas across
83 CONUS. By combining wavelet-based spectral analysis with a machine-learning classi-
84 fier trained on synthetic land surface model simulations, the proposed framework enables
85 irrigation detection directly from satellite soil moisture time series and reveals spatial
86 patterns of irrigated areas that may differ from those represented in existing satellite-
87 and inventory-based irrigation datasets, which often describe irrigation conditions for
88 a single reference year.

89 **2 Study area, methods and datasets**

90 The analysis was conducted over CONUS, which spans a wide range of hydrocli-
91 matic regimes and irrigation intensities, providing an ideal testbed for irrigation detec-
92 tion methods. We used the ESA SMOS Level-2 soil moisture product (v700; European
93 Space Agency (ESA) 2021) for the period 2010–2023. The product has an effective spa-
94 tial resolution of approximately 43 km and is geolocated on the equal-area Icosahedral
95 Snyder Equal Area (ISEA) 4H9 grid, with retrievals typically available every two to three
96 days. Recommended quality flags were applied to filter out radio-frequency interference,
97 poor retrievals, non-agricultural land, mountainous terrain, coastal areas, open water,
98 and urban surfaces (see Section S1.1 of the Supplementary Information, SI). The SMOS
99 observations were regridded to a regular 0.25° latitude–longitude grid (EPSG:4326) us-
100 ing nearest-neighbour interpolation and aggregated to weekly resolution. This grid is aligned
101 with the Noah–MP model grid used in the simulations.

102 The Landsat-derived Global Rainfed and Irrigated Cropland product (LGRIP30;
103 Teluguntla et al., 2023), shown in Figure S1 of the SI for the nominal year 2015, was used
104 as a reference dataset to develop the method and evaluate the SMOS-based estimates
105 of irrigated area. The original LGRIP30 maps have a spatial resolution of 30 m. To match
106 the spatial resolution of the model simulations and SMOS observations, the dataset was
107 aggregated to the same 0.25° (see SI section 1.3 for further details). Annual state-level
108 irrigation volumes from the USDA Farm and Ranch Irrigation Survey (FRIS; Section
109 S1.3 of the SI) for the years 2013 and 2018 were used as benchmark data for the Noah–MP
110 simulations of irrigation amounts.

111 Soil moisture estimates were obtained with the Noah–MP land surface model (v4.0.1;
 112 Niu et al., 2011) implemented in the NASA Land Information System (S. V. Kumar et
 113 al., 2006; Peters-Lidard et al., 2007). The model was run on the same regular 0.25° grid
 114 as the reprocessed SMOS data and forced with ERA5 forecast fields (Hersbach et al.,
 115 2020). Two deterministic Noah–MP simulations were performed with and without the
 116 dynamic sprinkler irrigation module (Ozdogan et al., 2010) (hereafter irrigation-on and
 117 irrigation-off simulations) using static maps of irrigated fraction from LGRIP30 and cli-
 118 matological green vegetation fraction seasonality. These simulations provide synthetic
 119 soil-moisture time series used for the wavelet analysis and for training the irrigation de-
 120 tection method.

121 The irrigation detection model developed in this study was then applied to SMOS
 122 observations to derive a satellite-based map of irrigated areas. This SMOS-derived ir-
 123 rigation map was subsequently used to constrain Noah–MP ensemble simulations with
 124 perturbed meteorological forcings to estimate irrigation volumes. A parallel ensemble
 125 simulation forced with the original LGRIP30 irrigation map was also performed for com-
 126 parison. A detailed description of the Noah–MP model setup, irrigation scheme param-
 127 eters, and simulation experiments is provided in the Supplementary Information (Sec-
 128 tions S1.4–S1.5).

129 We applied a Continuous Wavelet Transform (CWT; Section S1.6) using a Mor-
 130 let mother wavelet (Torrence & Compo, 1998) to decompose weekly soil-moisture time
 131 series into time–frequency space and compute global wavelet power spectra for Noah–MP
 132 irrigation-on and irrigation-off simulations and for SMOS observations. Wavelet differ-
 133 ences were computed between Noah–MP irrigation-on and irrigation-off soil moisture time
 134 series to train the classification method, and between Noah–MP irrigation-off simulations
 135 and SMOS soil moisture observations (Section S1.6.2) to detect irrigation under real-world
 136 conditions. Two metrics were derived from the wavelet analysis (Section S1.6.5): (i) the
 137 bias in the global wavelet power spectrum (M_p), computed as the difference between an
 138 irrigated signal (either Noah–MP irrigation-on or SMOS observations over a candidate
 139 irrigated area) and the corresponding Noah–MP irrigation-off simulation, averaged over
 140 the 6–18-month period band; and (ii) the reduction in wavelet coherence (M_C), defined
 141 as $M_C = 1 - \overline{R^2}_{BC}$, where $\overline{R^2}_{BC}$ is the wavelet coherence averaged over the same 6–
 142 18-month band between the two signals. The 6–18-month period range was selected to
 143 capture irrigation-driven modifications of the seasonal soil-moisture cycle while minimiz-

144 ing sensitivity to high-frequency retrieval noise at shorter periods and edge effects and
145 low-frequency hydroclimatic variability at longer periods.

146 A Random Forest classifier (RFC; Breiman, 2001) was trained on wavelet-derived
147 spectral features from simulated irrigation-on and irrigation-off soil moisture time series
148 to classify pixels as irrigated or non-irrigated (Section S1.7 of the SI). The LGRIP30 ir-
149 rrigation map was used as the target variable during training. Because the classifier is
150 trained only on spectral features derived from synthetic simulations and does not use ge-
151 ographic information, it cannot learn location-specific irrigation patterns from LGRIP30,
152 ensuring that its application is independent of location. To construct the training dataset,
153 additional predictors and perturbations were introduced. For each grid cell we computed
154 the aridity index (AI), defined as the ratio of potential evapotranspiration to precipi-
155 tation from ERA5 forecast fields. Binary irrigation labels were assigned using thresh-
156 olds of LGRIP30 irrigated fraction (10%, 20%, 30%, and 40% of total area). To account
157 for discrepancies between satellite observations and model simulations when applying
158 the classifier to SMOS data, Gaussian noise with zero mean and standard deviations be-
159 tween 5% (representing typical satellite soil-moisture uncertainty; Karthikeyan et al., 2017)
160 and 30% (worst-case scenario) was added to the simulated soil-moisture time series. These
161 perturbations represent the combined uncertainty arising from retrieval noise, model struc-
162 tural errors, and parameter uncertainty and, in general, do not reflect the accuracy of
163 the SMOS observations. The trained RFC was then applied to spectral differences be-
164 tween SMOS soil moisture observations and irrigation-off Noah–MP simulations.

165 The Receiver Operating Characteristic Curve (ROC) and the Area Under the Curve
166 (AUC) were used to assess the classification performance (Fawcett, 2006; Hanley & Mc-
167 Neil, 1982) for different thresholds on the irrigated area fraction and different noise lev-
168 els applied to the simulated irrigated soil-moisture time series, using LGRIP30 as a com-
169 parison dataset. Although LGRIP30 does not represent ground truth, it provides a re-
170 alistic, high-resolution reference for evaluating the similarity of the spatial patterns iden-
171 tified by our approach. The SMOS-derived binary irrigation maps were then scaled us-
172 ing the LGRIP30 irrigation fractions to constrain Noah–MP simulations of irrigation amounts.
173 This scaling is necessary because the SMOS classification determines the binary spatial
174 pattern of irrigated areas, while the LGRIP30 irrigation fraction constrains the total ir-
175 rrigated area within each 0.25° grid cell, thereby preserving the SMOS-derived spatial dis-

176 tribution. The resulting irrigation simulations (SMOS-derived) and simulations using
177 the original LGRIP30 irrigation map were evaluated against FRIS data.

178 Additional methodological details are provided in the SI, and a schematic overview
179 of the workflow is shown in Figure 1.

180 **3 Results**

181 To illustrate the expected differences in wavelet spectra between irrigation-on and
182 irrigation-off soil moisture time series, examples are shown for locations without and with
183 irrigation. In non-irrigated regions, the wavelet power spectra of SMOS and Noah-MP
184 (irrigation-off) are characterized by a dominant annual peak, and the bias near the 12-
185 month period approaches zero, indicating similar spectral content typical of rain-fed sys-
186 tems (Figure 2a; Figure S2). In contrast, irrigated regions exhibit a reduced contrast be-
187 tween the annual and semiannual peaks, with semiannual power approaching the mag-
188 nitude of the annual cycle (Figure 2b; Figure S3). This pattern reflects additional wet-
189 ting events associated with irrigation that are superimposed on the seasonal precipita-
190 tion cycle. This spectral shift is accompanied by a marked reduction in wavelet coher-
191 ence between SMOS and the irrigation-off Noah-MP simulation at 6–18-month periods
192 (Figure 2d), compared to non-irrigated areas (Figure 2c). The loss of coherence indicates
193 that irrigation introduces variability not captured by the model when irrigation is dis-
194 abled. The selected range of scales minimizes the influence of high-frequency noise in the
195 satellite retrievals and edge effects affecting both weekly variability and periods longer
196 than approximately two years. Overall, the spectral approach offers a clear advantage
197 over direct inspection of raw soil moisture time series by isolating irrigation-related low-
198 frequency variability from high-frequency observational noise, enabling a more robust
199 discrimination between irrigated and non-irrigated systems (see Figure S3b).

200 While the bias and coherence-drop metrics can be inspected at individual locations,
201 mapping irrigation consistently across CONUS requires a generalized framework. The
202 wavelet-based metrics derived from Noah-MP irrigation-on and irrigation-off simulations
203 were therefore used to train the RFC to distinguish irrigated from non-irrigated condi-
204 tions (see Methods). The spatial distributions of the predictor variables used by the clas-
205 sifier, including aridity index (AI), wavelet power bias, and coherence reduction (M_C),
206 are shown in Figure S4 of the SI. A clear pattern of increasing wavelet power bias is ob-

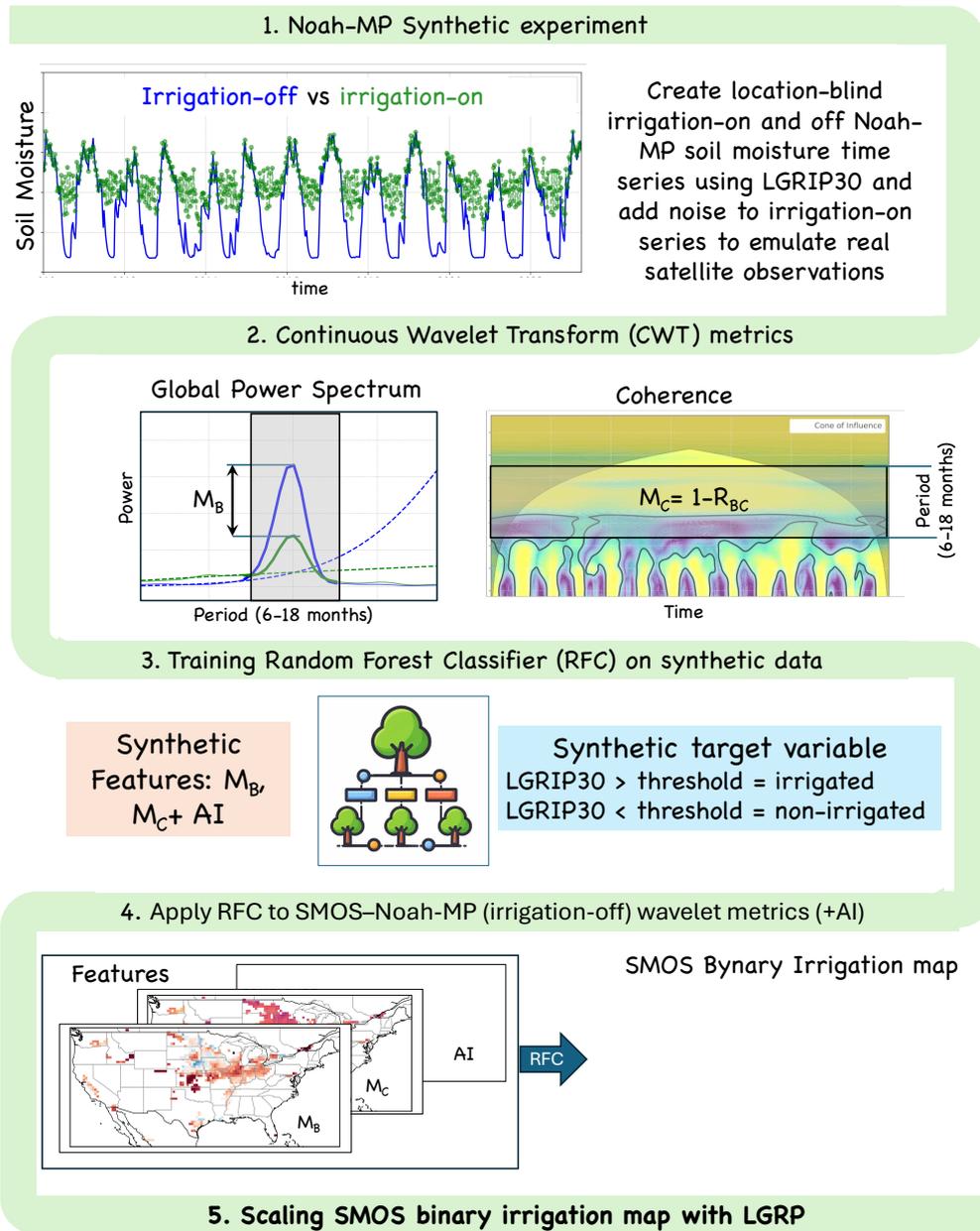


Figure 1. Schematic overview of the wavelet-based irrigation detection framework. (1) Synthetic soil-moisture (SM) time series are generated using the Noah-MP land surface model with irrigation-on and irrigation-off based on irrigation fractions from the Landsat-derived Global Rainfed and Irrigated Cropland product (LGRIP30). Multiple noise levels are added to the synthetic time series to emulate combined satellite retrieval uncertainty and model-data discrepancies. (2) A Continuous Wavelet Transform (CWT) is applied to paired soil moisture time series to compute the global wavelet power spectrum and wavelet coherence. From these spectra, irrigation-sensitive metrics are derived over the 6–18 month period range, including the bias in wavelet power and the reduction in wavelet coherence relative to simulations without irrigation. (3) A Random Forest classifier (RFC) is trained on synthetic experiments using aridity index (AI), wavelet power bias (M_B), and coherence reduction (M_C) as predictor variables and binary irrigation labels derived from LGRIP30 as targets (Section SI S1.6.5). (4) The trained RFC is applied to CWT-derived features from Soil Moisture and Ocean Salinity (SMOS) satellite observations and Noah-MP simulations irrigation-off to produce maps of irrigated and non-irrigated areas, (5) the binary classification map is scaled by LGRP. Wavelet spectra and coherence panels are shown schematically for illustration.

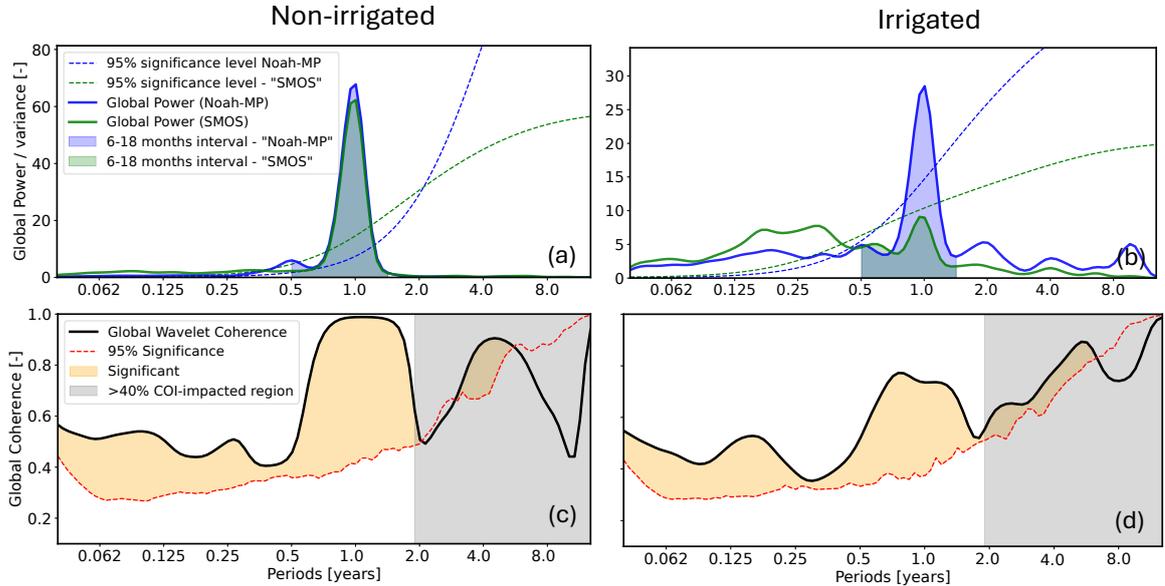


Figure 2. (a,b) Global wavelet power spectra of surface soil moisture from Noah-MP (irrigation-off) and SMOS for (a) a non-irrigated pixel in Eastern Washington (irrigation fraction $< 10\%$ according to LGRIP30) and (b) an irrigated pixel in South Dakota (irrigation fraction $\approx 80\%$). Dashed lines indicate the 95% significance levels relative to a red-noise background spectrum, following Torrence & Compo (1998). Thick lines highlight the 6–18-month period range used for the derivation of wavelet-based irrigation metrics. (c,d) Corresponding global wavelet coherence between Noah-MP (irrigation off) and SMOS for the same non-irrigated (c) and irrigated (d) pixels. The red dashed line denotes the 95% significance level based on Monte Carlo simulations assuming red noise (Torrence & Compo, 1998; Grinsted et al., 2004). Gray shading marks periods where more than 40% of the signal is affected by the cone of influence (COI).

207 served over intensively irrigated regions, along with a general north–south gradient in
208 M_C , indicating reduced coherence over irrigated regions.

209 The SMOS-derived irrigation map obtained using a 20% irrigation-fraction thresh-
210 old and assuming a noise level of 20% is shown in Figure 3. In this configuration, the
211 SMOS classification determines the spatial pattern of irrigated areas, while the LGRIP30
212 dataset constrains the irrigated fraction within each grid cell. Comparison with the LGRIP30
213 product (Figure S1) indicates that the major irrigated regions of CONUS, including Cal-
214 ifornia’s Central Valley, Washington State, the central–southern United States, and the
215 Great Plains, are well captured by the wavelet-based approach. Local discrepancies are
216 observed in areas with lower irrigated fractions, particularly within portions of the Great
217 Plains, where irrigation signals are weaker. Agreement between the two datasets also de-
218 creases in colder climates, where the seasonal overlap between precipitation and irriga-
219 tion reduces the distinctiveness of the spectral signature, and in the Corn Belt, where
220 dense vegetation may degrade the SMOS soil-moisture retrievals. Similar patterns of agree-
221 ment and discrepancy over CONUS have been reported in previous studies (Zaussinger
222 et al., 2019; S. Kumar et al., 2015; Crow et al., 2025).

223 A quantitative comparison with the LGRIP30 reference, assuming irrigated pix-
224 els where $LGRIP30 \geq 20\%$, yields an AUC of 0.7, indicating moderate agreement be-
225 tween the SMOS-based classification and the binary LGRIP30 irrigation map. The cor-
226 responding confusion matrix and the spatial distribution of true positives, true negatives,
227 false positives, and false negatives are provided in the SI (Table S2 and Figures S6). Sen-
228 sitivity analyses exploring alternative irrigation-fraction thresholds and noise assump-
229 tions are also reported in the SI Figure S7. These analyses show that lower noise lev-
230 els and smaller irrigation thresholds generally lead to reduced agreement (Section S2.1).

231 To assess whether the SMOS-derived irrigation map improves the representation
232 of irrigation volumes across CONUS, we compared Noah–MP irrigation estimates ob-
233 tained using the original LGRIP30 irrigation fractions (Figure S8) with those derived
234 from the SMOS-based irrigation map scaled by LGRIP30 within each 0.25° grid cell (Fig-
235 ure S8b). Simulated irrigation volumes were aggregated at the state level and evaluated
236 against irrigation statistics from the USDA Farm and Ranch Irrigation Survey (FRIS;
237 see SI Section S1.3).

238 Results (Figure S9) show consistently higher Pearson correlation coefficients when
 239 using the SMOS-based configuration than when using LGRIP30 directly. Correlation co-
 240 efficients between country-aggregated irrigation levels from FRIS and our simulations
 241 are $R = 0.92$ in both years with the SMOS-based approach, compared with $R = 0.81$
 242 in 2013 and $R = 0.78$ in 2018 for the LGRIP30-based configuration. Bias magnitudes
 243 are comparable between the two approaches but with opposite signs, with mean bias val-
 244 ues of $-2.05 \text{ km}^3 \text{ yr}^{-1}$ for the SMOS-based configuration and $2.07 \text{ km}^3, \text{ yr}^{-1}$ for the LGRIP30-
 245 based configuration, while RMSD values are similar (4.55 and $4.58 \text{ km}^3 \text{ yr}^{-1}$, respectively).
 246 These results indicate that incorporating the SMOS-based irrigation classification im-
 247 proves the representation of irrigation volumes in Noah–MP simulations. Additional de-
 248 tails of this analysis and the irrigation maps used in the simulations are provided in Sec-
 249 tion S2.2 of the SI.

250 Conclusions

251 We presented a new method to detect irrigation based on the spectral analysis of
 252 soil moisture time series. Using wavelet decomposition, we isolated irrigation-driven vari-
 253 ability at sub-annual scales and quantified two key metrics: the bias in wavelet power
 254 spectra and the reduction in wavelet coherence between SMOS observations and irrigation-
 255 off Noah–MP simulations.

256 By training a Random Forest classifier on synthetic soil-moisture time series sim-
 257 ulated with and without irrigation using Noah–MP, and by introducing different noise
 258 levels in the irrigation-on soil moisture signals and applying multiple thresholds to the
 259 LGRIP30 irrigation fractions to define binary irrigation labels, we generalized the rela-
 260 tionship between wavelet-based metrics, AI , and irrigation presence. Applied to SMOS
 261 observations and irrigation-off Noah–MP simulations over CONUS, the method success-
 262 fully mapped major irrigated regions, including California’s Central Valley, the High Plains,
 263 and parts of the southeastern United States. Comparison against the LGRIP30 prod-
 264 uct yielded an AUC of 0.7, demonstrating moderate-to-good discrimination despite model
 265 and retrieval uncertainties. Additionally, Noah–MP simulations constrained by the new
 266 SMOS-derived irrigation map improve the spatial representation of state-level irrigation
 267 water use across CONUS. The SMOS-based classification therefore provides an alterna-
 268 tive irrigation map derived from multi-year SMOS soil moisture observations, comple-
 269 menting static reference datasets such as LGRIP30.

270 The approach relies on land surface model simulations without irrigation to pro-
271 vide a reference representation of rain-fed soil moisture dynamics. In this study Noah–MP
272 was used for this purpose, but the framework is not tied to a specific model and could
273 be applied with other land surface models capable of simulating soil moisture variabil-
274 ity under rain-fed conditions.

275 This approach offers a scalable framework for irrigation detection at continental
276 to global scales. The method is currently limited by the coarse resolution of the SMOS
277 data and primarily complements existing methods based on higher-resolution optical ob-
278 servations. Future work will focus on characterizing the timing and intensity of irriga-
279 tion by exploiting the full wavelet spectrum and integrating higher-resolution observa-
280 tions from current and upcoming radar missions such as ROSE-L and NISAR. More gen-
281 erally, the framework is flexible and could potentially be applied using other types of in-
282 put variables that act as proxies for irrigation signals, provided that they capture irrigation-
283 induced anomalies in land surface conditions.

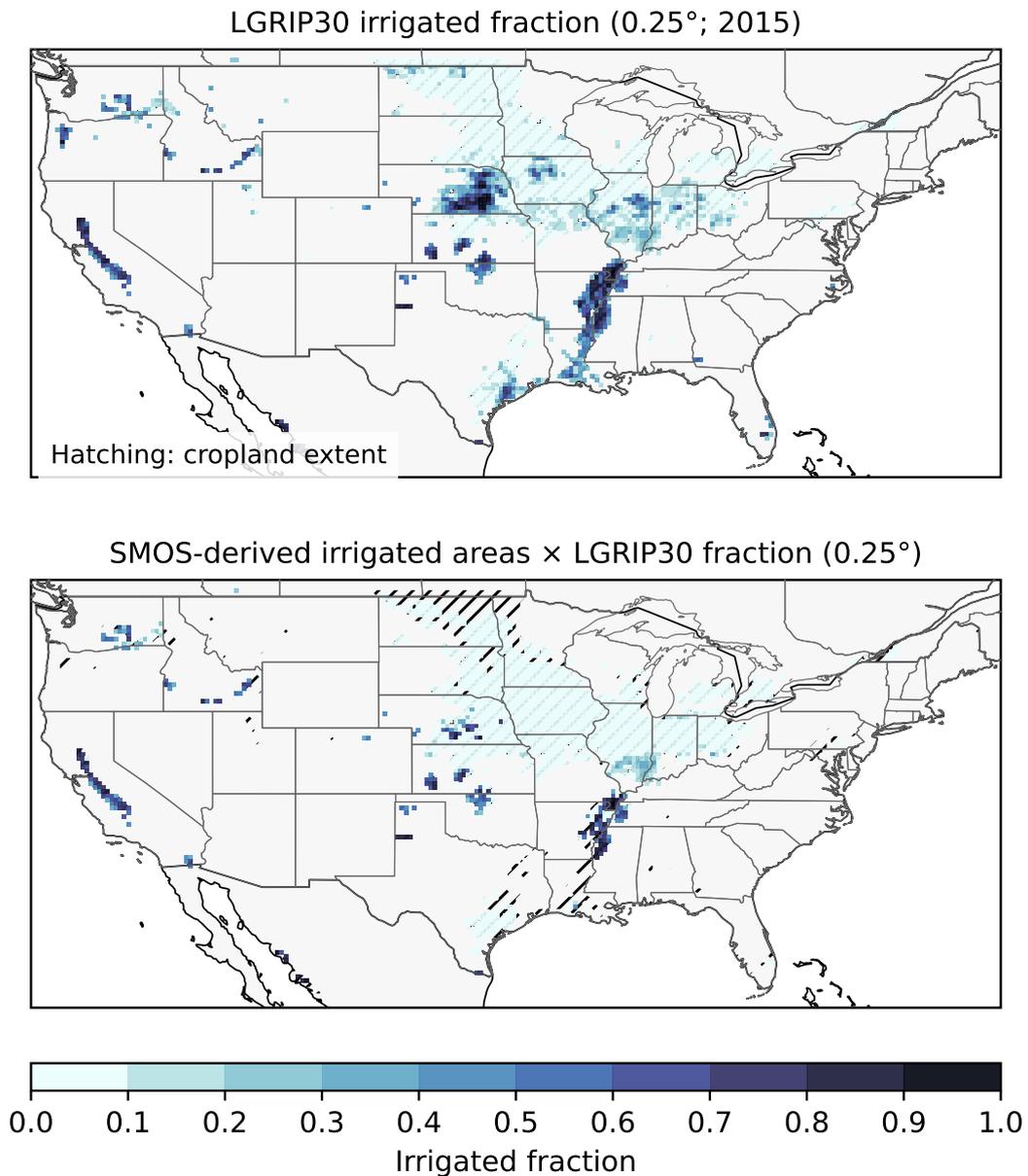


Figure 3. Irrigated fraction maps on the 0.25° grid over the Contiguous United States (CONUS). (a) Irrigated fraction from the Landsat-derived Global Rainfed and Irrigated Cropland product (LGRIP30; nominal year 2015) aggregated to the model grid. (b) Irrigated fraction obtained by combining the SMOS-based binary irrigation classification (20% threshold) with the LGRIP30 irrigated fraction within each grid cell. In this configuration, the SMOS classification determines the spatial pattern of irrigated areas, while the LGRIP30 fraction constrains the irrigated area within each 0.25° grid cell. Hatching indicates cropland extent. The color scale shows irrigated fraction divided into ten equal intervals between 0 and 1.

Open Research Section

SMOS Level-2 soil moisture data are publicly available from the European Space Agency (ESA) Earth Online portal (<https://earth.esa.int/eogateway>). ERA5 atmospheric forcing data were obtained from the Copernicus Climate Data Store (<https://cds.climate.copernicus.eu>). The Landsat-derived Global Rainfed and Irrigated Cropland product (LGRIP30) is publicly available from the NASA Land Processes Distributed Active Archive Center (LP DAAC). State-level irrigation statistics from the USDA Farm and Ranch Irrigation Survey (FRIS) are available from the U.S. Department of Agriculture (<https://www.nass.usda.gov>).

Noah-MP simulations were performed using the NASA Land Information System (LIS), which is publicly available at <https://lis.gsfc.nasa.gov>. The scripts used to perform the wavelet analysis and Random Forest classification, together with the derived irrigation classification maps and supporting diagnostics, will be deposited in a public repository (e.g., Zenodo), and a persistent DOI will be provided upon acceptance.

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