

1 Coastal winds, sea surface temperature, and upwelling in Atlantic 2 Canada in two years differing in ENSO and NAO conditions

3 **Ricardo A. Scrosati¹ & Julius A. Ellrich**

4 St. Francis Xavier University, Department of Biology, Antigonish, Nova Scotia B2G 2W5,
5 Canada

6 ¹ Corresponding author. Phone: +1-902-867-5289. Email: rscrosat@stfx.ca

7 **Abstract**

8 The most studied upwelling systems occur on eastern ocean boundary coasts. We hereby
9 focus on a western ocean boundary coast upwelling system located in Atlantic Canada. Using
10 daily in-situ data on sea surface temperature (SST), we demonstrate a marked contrast in cooling
11 between July 2014 (pronounced) and July 2015 (weak) for two locations ca. 110 km apart on the
12 southeastern coast of Nova Scotia. Southwesterly winds blow parallel to this coast. In July 2014,
13 such winds were common and their daily duration was negatively related to coastal SST. In July
14 2015, however, such winds were considerably less prevalent and unrelated to SST. These results
15 are consistent with a marked interannual contrast in wind-driven upwelling. Independent
16 measures of an upwelling index confirm this view, as the index was largely positive (indicative
17 of upwelling) during July 2014 but, albeit generally positive, often lower, and sometimes even
18 negative (indicative of downwelling), during July 2015. Both studied periods also experienced
19 considerable differences in climatic phenomena such as the El Niño Southern Oscillation
20 (ENSO) and the North Atlantic Oscillation (NAO). In July 2015, a strong El Niño event co-
21 occurred with a negative NAO index, while both phenomena exhibited near-zero indices in July
22 2014. Influences of these phenomena on coastal winds and upwelling are known for other
23 regions. Thus, given the relationships hereby reported for Atlantic Canada and the oceanographic
24 and ecological importance of upwelling, it would be profitable to further investigate the nature
25 and extent of atmospheric–oceanic links on this western ocean boundary coast.

26 **Introduction**

27 Coastal upwelling refers to the rise of deep, cool waters to the surface of nearshore pelagic
28 environments as warmer surface waters move offshore. As upwelled waters are often rich in
29 nutrients, they are typically associated to a high biological productivity and important fisheries
30 (FAO 2016). Several mechanisms can trigger coastal upwelling, but a common one involves the
31 combined action of winds and the Coriolis force. In the northern hemisphere, persistent
32 alongshore winds blowing with the coast on the left (on the right in the southern hemisphere)
33 generate an offshore surface Ekman transport that triggers coastal upwelling (Kämpf and
34 Chapman 2016).

35 There are four major wind-driven upwelling systems in the world, all of them located on
36 eastern ocean boundary coasts: the California Current system, the Humboldt Current system, the
37 Canary/Iberia Current system, and the Benguela Current system (Kämpf and Chapman 2016).
38 With generally a lower extent or prevalence, wind-driven upwelling also occurs on western
39 ocean boundary coasts, such as the coasts of Somalia (Currie et al. 1973) and Brazil (Mazzini
40 and Barth 2013).

41 Upwelling has also been reported for the Nova Scotia shore, another western ocean
42 boundary coast that is located in eastern Canada. Off southwestern Nova Scotia, upwelling is
43 mainly caused by tidal mixing and submarine topography (Tee et al. 1993; Chegini et al. 2018).
44 However, along the southeastern (open Atlantic) coast of Nova Scotia, upwelling is mainly
45 caused by wind action (Petrie et al. 1987; Shan et al. 2016). On this coast, upwelling is
46 particularly noticeable in July, favoured by southwesterly winds that blow parallel to the shore
47 (Petrie et al. 1987). However, recent information on interannual differences and comparisons
48 with other coasts are unavailable. To address this knowledge gap, the present paper examines
49 differences in July upwelling on this coast between 2014 and 2015, shows how coastal winds
50 relate to such variation, and provides values of an upwelling index to enable comparisons with
51 other coasts of the world. We also briefly discuss the relationship of the observed upwelling and
52 wind interannual changes with large-scale climatic phenomena.

53 **Materials and methods**

54 In July, seawater temperature in shallow (6–20 m deep) nearshore environments on central
55 and southern sections of the Atlantic coast of Nova Scotia decreases relative to June because of

56 wind-driven upwelling. August temperatures tend to return to the expected seasonal rise
57 according to heat flux from the atmosphere (Petrie et al. 1987). Thus, given our research goals,
58 we focused our study on the month of July, monitoring also late June as a reference. Specifically,
59 we measured sea surface temperature (SST) daily between 20 June and 31 July for two
60 consecutive years (2014 and 2015) at two intertidal locations ca. 110 km apart: Duck Reef
61 (44.4913, -63.5270) and Western Head (43.9896, -64.6607). Both locations consist of stable
62 bedrock, face the open ocean directly, and lack any human-made constructions nearby. At the
63 mid-to-high intertidal zone of each location, we used eyescrews and cable ties to permanently
64 attach submersible loggers (HOBO Pendant logger, Onset Computer, Bourne, MA, USA) to the
65 rocky substrate, allowing very little contact between the loggers and the substrate. At each
66 location, we installed four loggers in 2014 and two in 2015, replicate loggers being tens of m
67 apart. Each logger recorded temperature every 30 min. From the resulting time series, we
68 extracted the values of daily SST, here considered as the temperature recorded nearest to the time
69 of the highest tide of each day (when the loggers were submerged). We determined the time of
70 such tides using information for the closest tide reference stations to our locations: Sambro
71 (44.4833, -63.6000) for Duck Reef and Liverpool (44.0500, -64.7167) for Western Head (Tide
72 and Current Predictor 2018). For each location and year, SST was highly correlated between
73 replicate loggers (average $r = 0.98$), so we averaged the corresponding daily values to generate
74 one time series of daily SST data for each location and year. We then compared SST dynamics
75 between 2014 and 2015 for both locations.

76 Since SST dynamics suggested a stronger upwelling in 2014 than in 2015 (see Results), we
77 investigated if wind data could explain such a difference, under the knowledge that
78 southwesterly winds trigger upwelling on this coast (Petrie et al. 1987). The studied coast has an
79 orientation of ca. 60° measured clockwise from the north (0°). Thus, for each day of July 2014
80 and July 2015, we counted for each location the number of hours in which winds blew from the
81 southwest (hereafter, "duration of SW winds") from any direction between 230° – 250° measured
82 clockwise from the north. We made these calculations using data for the closest wind reference
83 stations to our locations: Herring Cove Buoy (44.5586, -63.5453; SmartAtlantic 2018) for Duck
84 Reef and Western Head Station (43.9900, -64.6642; Government of Canada 2018) for Western
85 Head. We then calculated the Pearson correlation coefficient between the duration of SW winds
86 and daily SST for each location and year.

87 Finally, for each day of July 2014 and July 2015 for each location, we calculated a
88 commonly used upwelling index (UI) that is based on wind data and geographic information.
89 First, we calculated wind stress (equation 4.2 in Stewart 2008) by multiplying air density (1.28
90 kg m^{-3}), the wind drag coefficient (0.0015; Kämpf and Chapman 2016), and the square of wind
91 speed (in m s^{-1}). Wind speed data were available hourly for the two wind reference stations
92 mentioned above, so we calculated wind stress for each hour. Then, we calculated Ekman
93 transport (equation 2.2 in Kämpf and Chapman 2016) by dividing wind stress by seawater
94 density (1026 kg m^{-3}) and by the Coriolis parameter, which depends on latitude. We calculated
95 UI (equation 2.3 in Kämpf and Chapman 2016) by multiplying Ekman transport by the cosine of
96 an angle (α) that represented the difference between the orientation of our coast (60°) and the
97 angle (also measured clockwise relative to the north) denoting wind direction with the wind
98 vector's origin centered in the target location. For example, for our coast, a southwesterly wind
99 blowing perfectly parallel to the shore would yield $\alpha = 0^\circ$. Wind direction data were also
100 available hourly for the two wind reference stations mentioned above. To calculate daily UI
101 values, we averaged the corresponding hourly UI values. In this paper, we express UI values in
102 terms of cubic meters of seawater transported per second per 100 m of coastline. Positive UI
103 values indicate upwelling, whereas negative values indicate downwelling (Menge and Menge
104 2013). The dataset used for the analyses described in the Results section is available online
105 (Scrosati and Ellrich 2018a).

106 **Results**

107 In late June 2014, coastal SST ranged between 9-14 °C at Duck Reef and 8-13 °C at
108 Western Head. However, during the first half of July, SST exhibited a persistent decrease at both
109 locations, reaching minima of 6 °C at Duck Reef and 4 °C at Western Head (Fig. 1). The rest of
110 July showed frequent alternations of drops and rises of SST at both locations (Fig. 1). In late
111 June 2015, the SST range was similar to that recorded in the same period of the previous year.
112 However, in July 2015, SST did not exhibit the persistent decrease seen in July of the previous
113 year, in fact showing a mildly increasing trend (also with some alternations between drops and
114 rises) towards the end of the month (Fig. 1).

115 Southwesterly winds were more common in 2014 than in 2015. The mean daily duration of
116 SW winds was 10 h (range = 0-22 h) in July 2014 and only 2 h (0-9 h) in July 2015 at Duck Reef

117 and 6 h (0-12 h) in July 2014 and only 3 h (0-9 h) in July 2015 at Western Head. Daily SST was
 118 negatively related to the daily duration of SW winds in July 2014 ($r = -0.43$, $p = 0.016$ for Duck
 119 Reef and $r = -0.38$, $p = 0.037$ for Western Head), but both variables were uncorrelated in July
 120 2015 ($r = 0.06$, $p = 0.749$ for Duck Reef and $r = -0.27$, $p = 0.141$ for Western Head; Fig. 2).

121 UI was overwhelmingly positive in July 2014, peaking at $190 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹ at
 122 Duck Reef and $161 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹ at Western Head. Both peaks occurred in the
 123 first week of July (Fig. 3), just a few days before the lowest SST values were reached (Fig. 1).
 124 Similarly high UI values were uncommon, however, as 81 % of the 27 positive values of daily
 125 UI from Duck Reef were lower than $100 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹, while 86 % of the 28
 126 positive values of daily UI from Western Head were lower than $50 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹.
 127 The few negative UI values recorded in July 2014 were not lower than $-6 \text{ m}^3 \text{ s}^{-1}$ (100 m of
 128 coastline)⁻¹ (Fig. 3).

129 In July 2015, positive UI values were also predominant, but generally lower than in July
 130 2014, reaching peaks of only $68 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹ at Duck Reef and $52 \text{ m}^3 \text{ s}^{-1}$ (100 m
 131 of coastline)⁻¹ at Western Head (Fig. 3). Negative values occurred more frequently than in July
 132 2014, reaching minima of $-65 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹ at Duck Reef and $-16 \text{ m}^3 \text{ s}^{-1}$ (100 m of
 133 coastline)⁻¹ at Western Head (Fig. 3).

134 Overall, mean monthly UI (considering all 31 daily values) was $57 \text{ m}^3 \text{ s}^{-1}$ (100 m of
 135 coastline)⁻¹ at Duck Reef and $30 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹ at Western Head in July 2014 and
 136 $11 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹ at both Duck Reef and Western Head in July 2015.

137 **Discussion**

138 Past studies had determined that wind-driven upwelling occurs on the Atlantic coast of
 139 Nova Scotia, especially in July (Petrie et al. 1987; Shan et al. 2016). The present study has
 140 revealed that marked interannual variation can exist. This was evident in the dynamics of SST,
 141 as seawater cooling (relative to late June) was prevalent in July 2014 but inconspicuous in July
 142 2015. These results are valuable because they are based on in-situ SST data (as opposed to
 143 satellite data, which may be less accurate and sometimes even unavailable for some days).

144 The independent data on wind direction explained the basic interannual difference in SST
 145 dynamics under a model of wind-driven coastal upwelling. This is so because southwesterly

146 winds (which favour coastal upwelling and thus cooling; Petrie et al. 1987) were considerably
147 more common in July 2014 than in July 2015. Lastly, UI values, also obtained independently
148 from SST data, indicated that coastal upwelling was more frequent and intense in July 2014 than
149 in July 2015 (when a few days with clear downwelling conditions were even registered).

150 A recent study (Menge and Menge 2013) measured UI for various coastal locations in
151 California, Oregon, and New Zealand. Our highest values for Nova Scotia, $190 \text{ m}^3 \text{ s}^{-1}$ (100 m of
152 coastline)⁻¹ for Duck Reef and $161 \text{ m}^3 \text{ s}^{-1}$ (100 m of coastline)⁻¹ for Western Head (measured in
153 July 2014), are similar to the highest values published by Menge and Menge (2013) for
154 California, a region of intense summer upwelling. However, the published California values are
155 averages for several months and years, so values for peak days would be higher than ours. In
156 addition, as described in Results, high UI values were in fact uncommon in Nova Scotia. The
157 July means of UI reported in Results actually place Nova Scotia in the range of values that
158 typically describe locations in Oregon and western New Zealand, where upwelling is, on
159 average, less intense and more intermittent than in California (Menge and Menge 2013).

160 As winds are central to the occurrence of upwelling on the southeastern coast of Nova
161 Scotia (Petrie et al. 1987; Shan et al. 2016), an emerging question is what caused the interannual
162 wind differences described above. The El Niño Southern Oscillation (ENSO) is a climatic
163 phenomenon that, through alterations of wind patterns, greatly affects the air and sea along the
164 equatorial Pacific (Philander 1990). However, this phenomenon is also the largest source of year-
165 to-year fluctuations in global weather patterns (Stewart 2008; Timmermann et al. 2018).
166 Teleconnections between ENSO and weather patterns (including winds) in North America have
167 been identified (George and Wolfe 2009; Wu and Lin 2012; Whan and Zwiers 2017). Thus,
168 influences on the interannual wind changes hereby reported for Nova Scotia seem possible.
169 Interestingly, the persistent upwelling seen in July 2014 was associated not only to frequent
170 southwesterly winds but also to an absence of El Niño or La Niña conditions, as the Oceanic
171 Niño Index (ONI) centered on that month was 0.1 (NOAA 2018a). The upwelling reported by
172 Petrie et al. (1987) for July 1984 was also associated to a lack of El Niño or La Niña conditions,
173 as ONI was then -0.3 (NOAA 2018a). Conversely, the weak upwelling and infrequent
174 southwesterly winds in July 2015 occurred during a strong El Niño event, as ONI was then 1.5
175 (NOAA 2018a).

176 The North Atlantic Oscillation (NAO) is another large climatic phenomenon. It mainly
177 affects the North Atlantic basin, although it also influences atmospheric variability at middle and
178 high latitudes elsewhere in the northern hemisphere (Hurrell 1995; Whan and Zwiers 2017). In
179 relation to our findings, it is interesting to note that, like the ONI, the NAO Index differed greatly
180 between July 2014 (0.2) and July 2015 (-3.2; NOAA 2018b). Thus, as found for other weather
181 phenomena in the northern hemisphere (Wu and Lin 2012; Hafez 2017), both ENSO and NAO
182 might contribute to alter wind patterns on the Nova Scotia coast, thereby affecting coastal
183 upwelling (see Santos et al., 2005, for a NAO example elsewhere). The effects of ENSO and
184 NAO could even be synergistic when a high ONI and low NAO Index co-occur (Wu and Lin
185 2012), as was the case in July 2015.

186 Overall, the relationships reported in this paper suggest that atmospheric-oceanic coupling is
187 worth investigating further in Atlantic Canada. Such links are relevant not only environmentally
188 (Varela et al. 2018) but also ecologically, as upwelling is thought to influence pelagic primary
189 production (Menge and Menge 2013) with bottom-up effects on coastal benthic species, as recent
190 data suggest for Nova Scotia (Scrosati and Ellrich 2018b).

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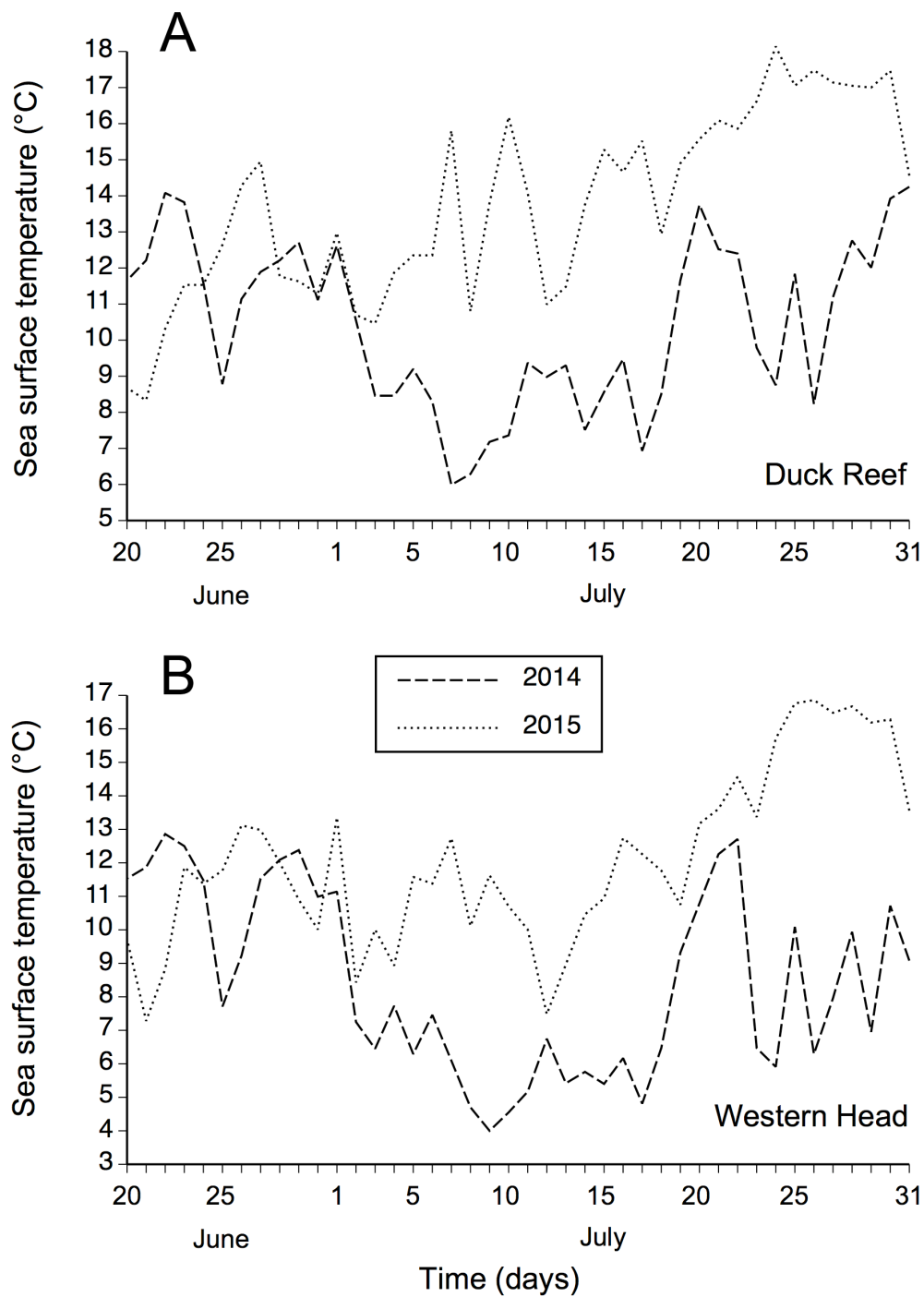
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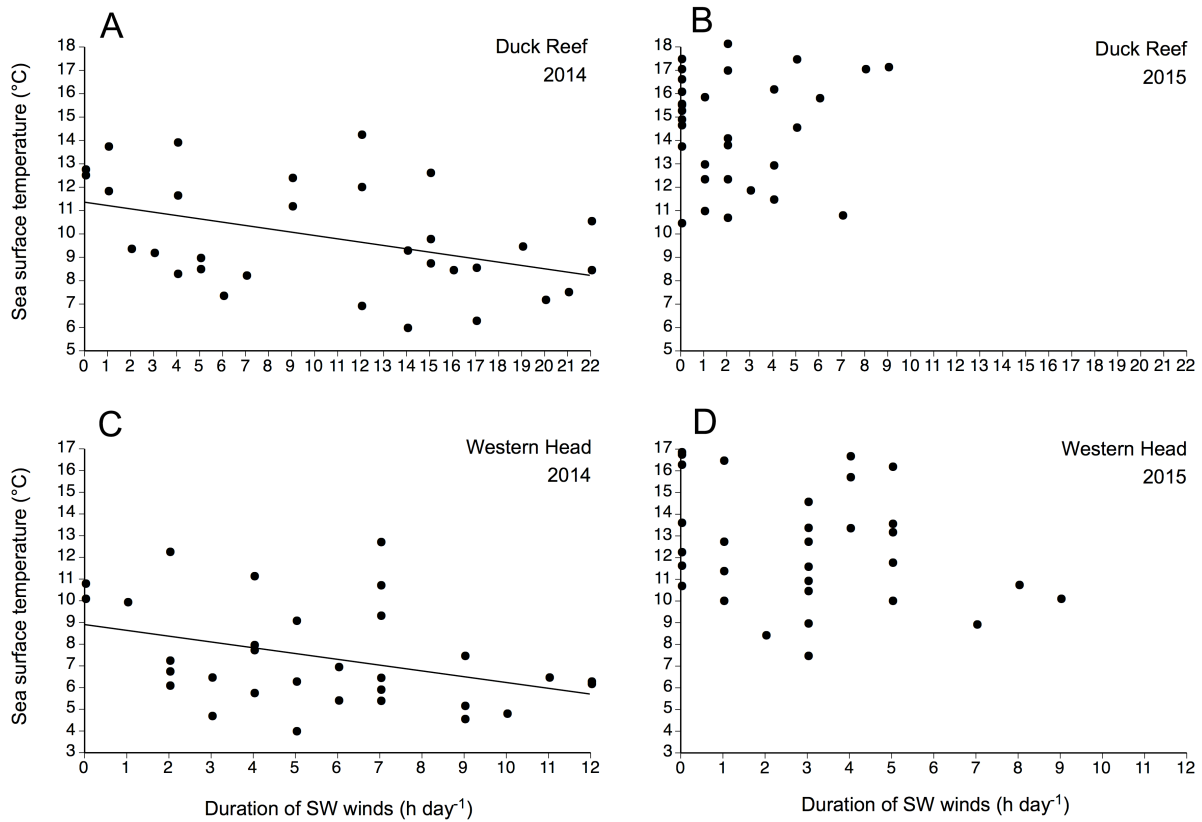
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258

259 **Fig. 1.** Daily change of sea surface temperature between 20 June and 31 July 2014 and 2015 at
 260 (A) Duck Reef and (B) Western Head.

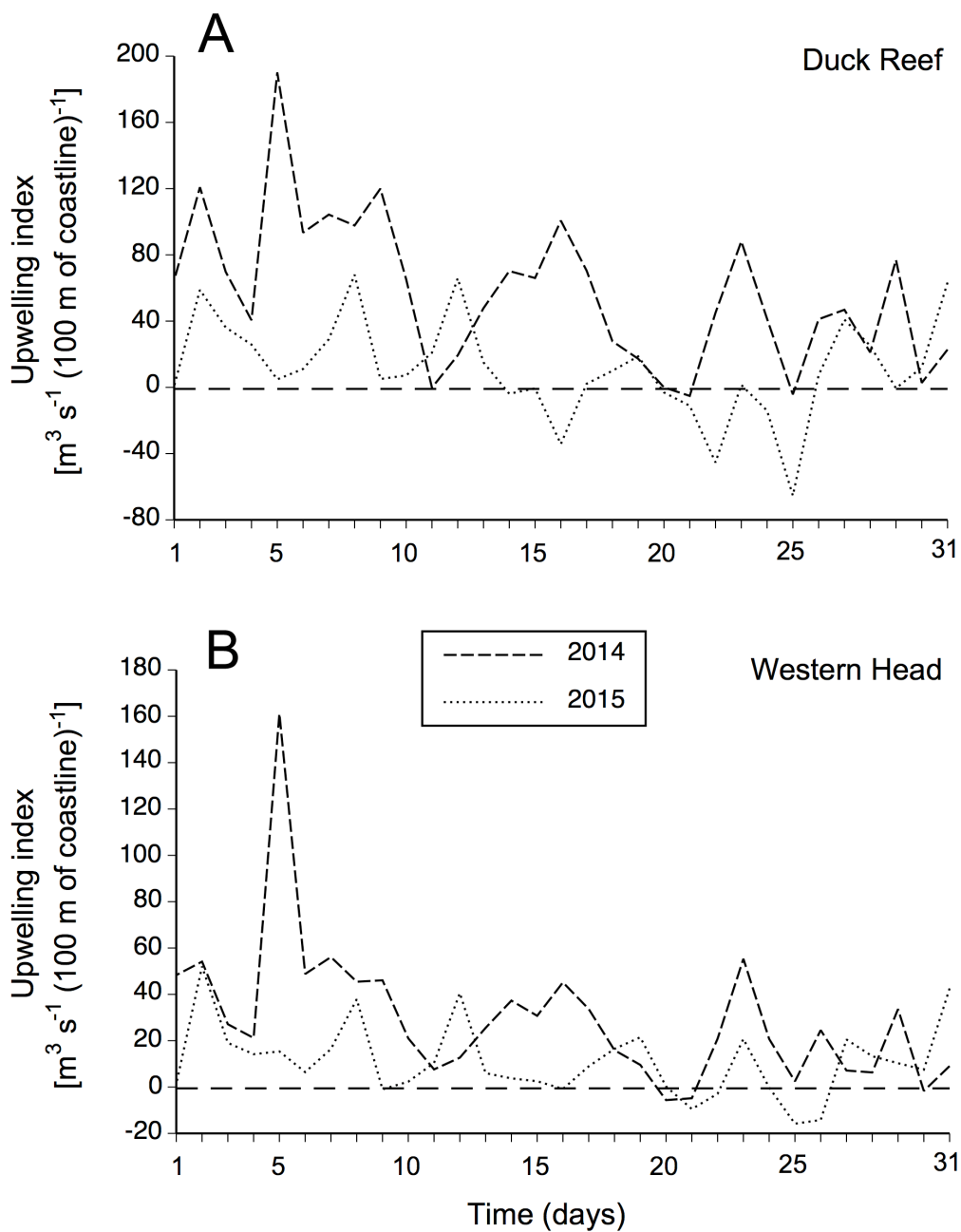
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261

262 **Fig. 2.** Daily sea surface temperature versus daily duration of SW winds for Duck Reef in (A)
 263 2014 and (B) 2015 and for Western Head in (C) 2014 and (D) 2015. Only (A) and (C) exhibited
 264 a significant relationship (see Results for details).

265



265

266 **Fig. 3.** Daily change of the upwelling index between 1–31 July 2014 and 2015 at (A) Duck Reef
 267 and (B) Western Head.