

1 A nearly 1.5 millennia long record of North Atlantic climatic forcing recorded in tree-ring growth
2 records from Angstel/Vecht Delta, the Netherlands (1283 BCE-156 CE)

3
4 Michiel Arts^{1*}

5 Jos Bazelmans^{2,3*}

6
7 *1 - SEDICLIM lab, Department of Geology, University of Liège, Belgium*

8 *2 - Cultural Heritage Agency of the Netherlands and Faculty of Archaeology*

9 *3 - University of Leiden, Leiden, Netherlands*

10 *These authors contributed equally to this work.

11
12 Corresponding Author Email: j.bazelmans@cultureelerfgoed.nl

13
14
15 This manuscript is a non-peer-reviewed preprint submitted to
16 EarthArXiv and is currently under review at The Holocene.
17

18
19 Keywords: tree-ring growth, bog oak, alder carr, Angstel and Vecht, Netherlands,
20 hydroclimate variability, North Atlantic Oscillation (NAO), Atlantic
21 Multidecadal Oscillation (AMO), bicoherence, groundwater dynamics,
22 (cross) wavelet analysis, climate forcing
23

24 Abstract

25

26 The impact of multi-annual to decadal climatic cycles on tree-ring growth remains
27 understudied. Here, we analyse a nearly one-and-a-half-thousand-year-long tree-ring
28 chronology (1283 BCE–156 CE) based on bog oaks from an alder carr environment in the
29 Angstel-Vecht delta in the Netherlands to assess the influence of multi-year climatic
30 oscillations on tree-ring growth and the origin of the identified cycles. Wavelet,
31 crosswavelet, and bicoherence spectral analyses reveal significant cycles ranging from 2.2 to
32 100 years. To investigate the forcing mechanisms behind these cycles, we compare the tree-
33 ring record with the Uamh an Tartair stalagmite record from northwestern Scotland,
34 reflecting large-scale variability in North Atlantic precipitation. Cross-wavelet analysis reveals
35 a quarter-cycle phase lag across multiple timescales, indicating a shared North Atlantic
36 climatic driver, with the stalagmite responding almost instantaneously, whereas the
37 Angstel/Vecht tree-ring record shows a delayed response. This delay is attributed to
38 cumulative climate forcing mediated by groundwater changes and hydrological storage
39 processes within the peatland system. The phase relationships suggest that, alongside
40 flooding events, periods of increased regional rainfall elevate groundwater levels,
41 suppressing oak growth in the Angstel/Vecht delta. The 100-, 60-, and 24-year cycles
42 correspond to known climatic oscillations linked to North Atlantic climate modes such as the
43 North Atlantic Oscillation (NAO) and the Atlantic Multidecadal Oscillation (AMO), or possibly
44 solar cycles. Some cycles, such as the 40- and 15-year cycles, cannot be directly attributed to
45 external forcing but instead appear to be nonlinear combination tones. These results
46 indicate that cyclicity in tree-ring growth reflects non-linear interactions among multiple
47 forcing mechanisms across timescales, and demonstrate that lowland wetland bog oak
48 records provide sensitive archives of multiannual to multidecadal hydroclimatic variability in
49 the North Atlantic region over long timescales.

50 1. Introduction

51

52 Hydroclimate variability in Western Europe is strongly influenced by large-scale climate
53 patterns over the North Atlantic, especially the North Atlantic Oscillation (NAO), the Atlantic
54 Multidecadal Oscillation (AMO), and solar cycles (Ait Brahim et al., 2018; Börgel et al., 2020;
55 Dieppois et al., 2013; Lüdecke et al., 2020; Pinto and Raible, 2012). These patterns affect
56 temperature, rainfall, storm activity, and groundwater recharge over seasonal to centennial
57 timescales, leaving clear traces in ocean, lake, and cave records (Ait Brahim et al., 2018;
58 Baldini et al., 2008; Denniston et al., 2008; Hubeny et al., 2006; Rust et al., 2020). However,
59 how these climate patterns are reflected in tree-ring data from lowland areas—especially
60 before the 19th century—remains unclear, making it difficult to identify cycles and
61 nonlinear interactions among different climate influences.

62

63 Tree rings provide records of annual hydroclimate variability. In the Angstel/Vecht region of
64 the central Netherlands, oaks grew in alder carrs on peat. Here, groundwater was closely
65 linked to precipitation and discharge from the Rhine River (Jansma, 2020). High rainfall and
66 runoff raised groundwater levels, prolonged anaerobic soil conditions, shortened the
67 growing season, induced root decay, impaired ectomycorrhizal functioning, and increased
68 the risk of oak mortality (Coder, 1994; Copini et al., 2016; Sass-Klaassen and Hanraets,
69 2006). Drier conditions allowed wider ring increments by improving root conditions (Sass-
70 Klaassen and Hanraets 2006). Both wetter and drier hydrology affect tree growth and
71 mortality. Prolonged high groundwater is suggested to have suppressed germination,
72 depressed growth, and increased mortality. Continued peat formation favoured the
73 preservation of subfossil wood by limiting microbial decay in waterlogged settings
74 (Leuschner et al., 2002; Sass-Klaassen and Hanraets, 2006).

75

76 Atmospheric pressure patterns associated with the positive NAO phase are closely linked to
77 major flood events in northwestern Europe (Brönnimann et al., 2025). The NAO index
78 expresses the pressure difference between the Azores High and the Icelandic Low. This
79 difference controls the strength and direction of westerly winds and storm tracks across the
80 North Atlantic (Domeisen et al., 2018; Eade et al., 2022; Folland et al., 2009; Hurrell et al.,
81 1997; Hurrell et al., 2003). During positive NAO phases, conditions are generally mild, wet,

82 and stormy. Negative phases bring colder, drier, and calmer conditions. These shifts are
83 particularly pronounced in winter. The NAO strongly influences winter and spring discharge
84 in the Rhine River (Steirou et al., 2017). Positive NAO phases correspond to elevated winter
85 discharge due to increased precipitation and earlier snow melt. Negative phases correspond
86 to reduced winter discharge followed by a (late) spring or early summer runoff peak.
87 Variability in the NAO occurs on multi-annual timescales. Multi-decadal to centennial
88 intervals of persistent NAO states are identified in paleoclimatic archives (Baker et al., 2015;
89 Faust et al., 2016; Olsen et al., 2012; Seip et al., 2019). Despite extensive research,
90 reconstructions of long-term NAO variability differ due to uncertainties in age models and
91 limitations in the calibration strength of NAO proxies (Baker et al., 2015).

92

93 To clarify how North Atlantic climate variability shapes terrestrial hydroclimate, we analyse
94 a 1.5 millennia-long tree-ring chronology (1283 BCE - 156 CE) from the Angstel/Vecht delta.
95 This chronology represents the most robustly replicated segment ($n \geq 10$) of a larger dataset
96 spanning 1371 BCE–488 CE (NLBOG26) with 90,914 measured rings. Our central goal is to
97 determine if this tree-ring record contains signatures of known atmospheric and solar
98 oscillations and their nonlinear interactions. We use the continuous wavelet transform
99 (CWT), Taner filtering, bicoherence spectral analysis, and the cross-wavelet transform (XWT)
100 to identify dominant periodicities and interactions in tree growth that may be linked to
101 major climate modes and solar forcing.

102

103 We use CWT to identify and track dominant cycles in the record. Based on these results, we
104 extract principal cycles using a Taner filter and CWT spectral power to evaluate their
105 persistence. Next, we apply bicoherence spectral analysis to determine the extent to which
106 these cycles are influenced by nonlinear interactions among frequency components. We
107 then isolate and compare subsets of nonlinear signals to those generated from higher-order
108 combinations of cycles, testing whether observed interactions produce new cyclic
109 components. Finally, we assess if the identified variability reflects large-scale North Atlantic
110 forcing by performing cross-wavelet analysis between the Angstel/Vecht record and the
111 Uamh an Tartair stalagmite record. This integrated approach directly quantifies how
112 Angstel/Vecht tree growth responds to climatic oscillations and determines how lowland

113 wetland tree-ring archives capture hydroclimatic variability linked to ocean-atmosphere
114 dynamics and solar forcing.

115

116 2. Paleogeographic setting

117

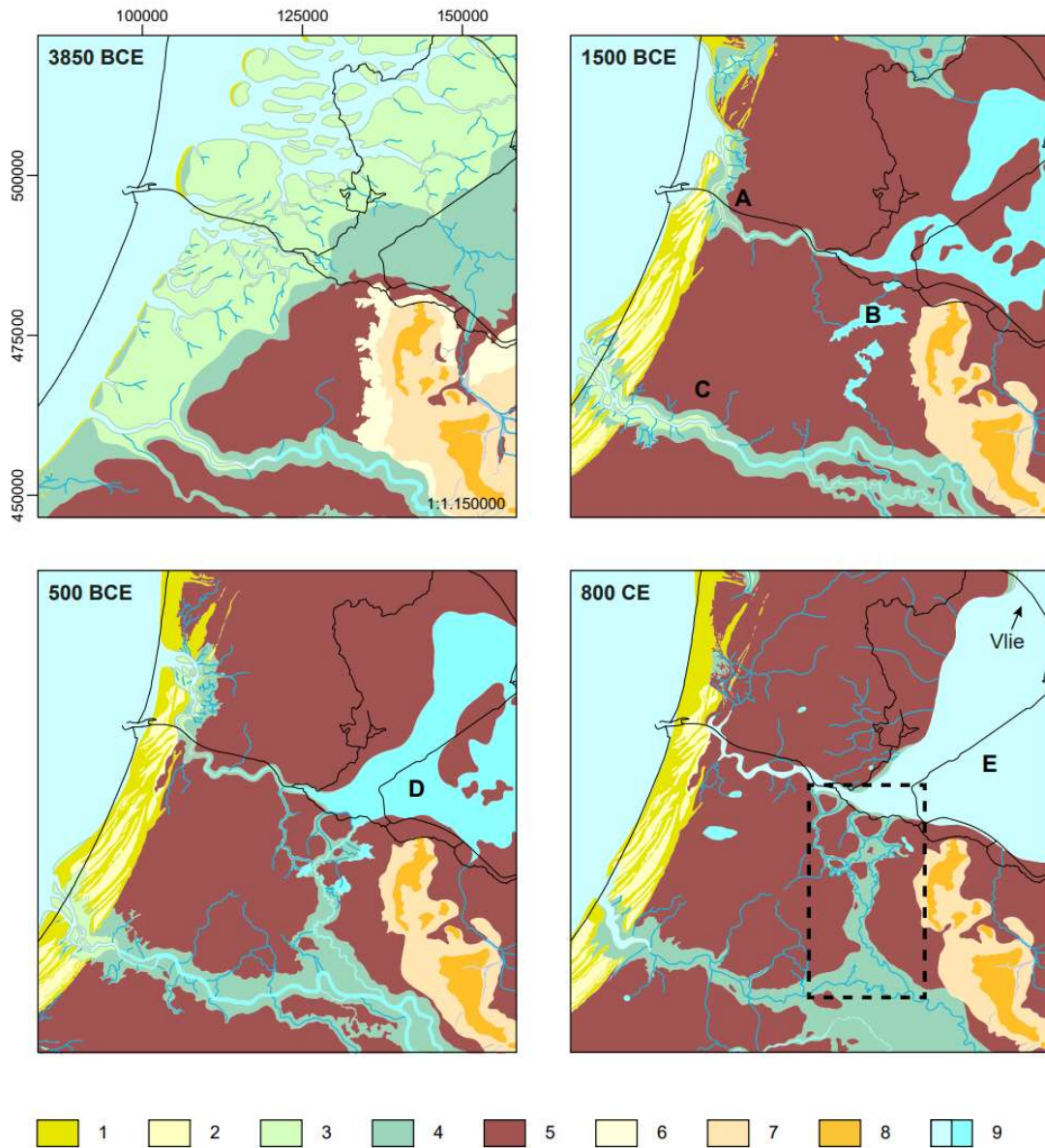
118 The Angstel/Vecht delta, located between the Dutch cities of Amsterdam, Utrecht, and
119 Naarden, and the former Zuiderzee/Lake Flevo, represents a Holocene peat-dominated area
120 at the northern margin of the Rhine delta (Van Asselen et al., 2017; Vos et al., 2020; Weerts
121 et al., 2002)(Figure 1). The area developed on an undulating Late Weichselian coversand
122 surface that became buried beneath Holocene peat, lacustrine and fluvial deposits as
123 groundwater levels rose in response to post-glacial sea-level rise, promoting extensive peat
124 formation across the coastal plain (Bos et al., 2009, 2012; Hijma et al., 2019).

125

126 The Angstel/Vecht system developed as a distributary of the Old Rhine from the 12th
127 century BCE onwards. Until the early Middle Ages, the Angstel and Vecht rivers conveyed up
128 to about 10% of the Rhine discharge to the North Sea. This formed a dynamic flood basin
129 with interbedded fluvial, lacustrine, and peat deposits (Van der Woude, 1984; Vos et al.,
130 2020; Weerts et al., 2002). Progressive infilling of peat lakes was followed by the
131 development of levees through repeated overbank deposition of clay-rich sediments. This
132 influenced the spatial distribution of peat facies, with wood peat forming in river-influenced
133 areas and more distal zones characterised by rain-fed peat accumulation (Brouwer et al.,
134 2002; De Lange and Bles, 1963; Poelman, 1966; Van Dinter et al., 2017). These processes
135 further structured the landscape and controlled local drainage patterns (Bos et al., 2009).

136

137 Thick peat sequences accumulated across this wetland, supporting alder carrs and scattered
138 oaks. These oaks were sensitive to fluctuations in groundwater levels (Jansma 2020; Sass-
139 Klaassen and Hanraets 2006). Elevated water reduced tree-ring growth by limiting root
140 aeration. Drier phases enabled wider ring increments. This hydrological dependence makes
141 the site ideal for reconstructing past hydroclimate. Subfossil oaks in the peat form the
142 Angstel/Vecht tree-ring chronology NLBOG26, detailed by Bazelmans et al. (submitted). In
143 this setting, the Angstel/Vecht region is a highly sensitive archive of externally forced
144 hydroclimatic variability.



146

147

148 Figure 1. Caption : Figure 1. Palaeogeography of the Angstel-Vecht delta during the 3850

149 BCE, 1500 BCE, 500 BCE and 800 CE time slices. A, Oer-IJ; B, Lake Aetsveld; C, Old Rhine; D,

150 Lake Flevo; E, Almere. 1, beach barriers and low dunes; 2, beach planes and dune valleys; 3,

151 tidal flats; 4, salt marshes, floodplains and stream valleys; 5, peat areas, 6, Pleistocene sand

152 areas below modern sea level, 7, Pleistocene sand areas above modern sea level; 8, ice-

153 pushed ridges; 9, outer and inner water. Dashed line: research area. Modified after Vos et

154 al., (2018).

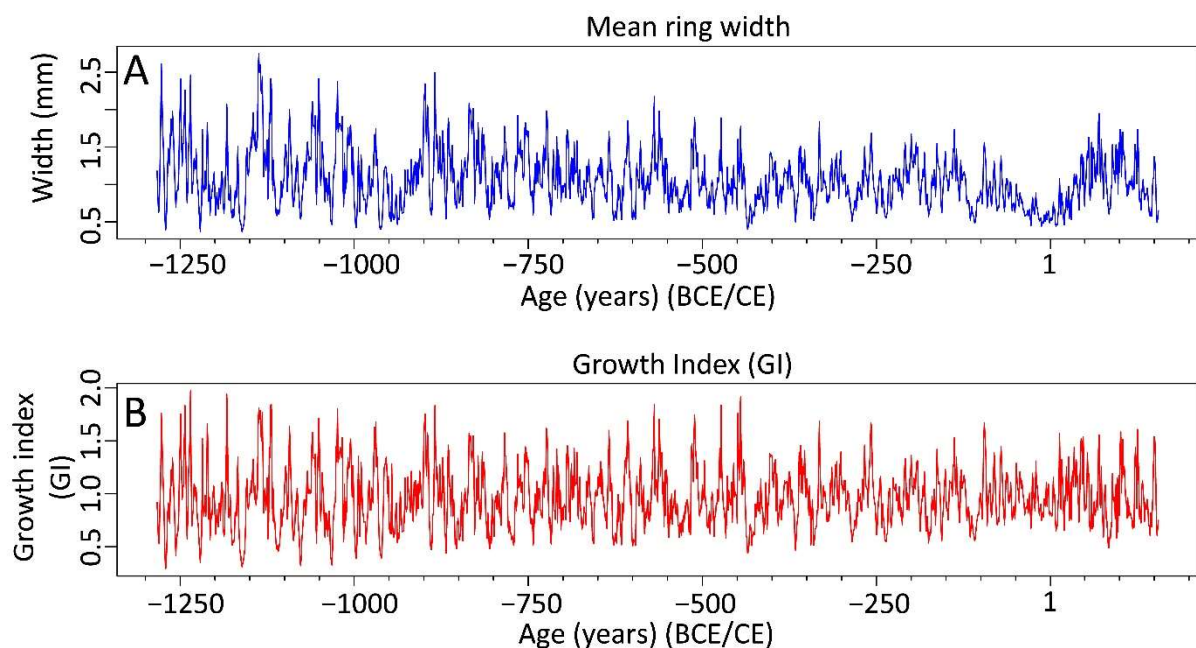
155

156 3. Materials and Methods

157

158 To uncover cyclicity in the Angstel/Vecht tree-ring chronology, we analysed the growth
159 index (GI) record NLBOG26 (Bazelmans et al., submitted; Figure 2, upper panel). This record
160 contains 526 specimens from 31 sites; 500 are absolutely dated and together form a robust
161 chronology for the Angstel/Vecht delta. Raw ring-width measurements include long-term
162 biological trends that can obscure climatic signals. To address this, we used the detrended
163 GI record. We removed age-related growth trends using standard detrending, including a
164 128-year spline with a 50% variance cut-off (Cook et al., 1990). The calculation was tailored
165 to each time series, using a spline length equal to 67% of the series length, with a maximum
166 spline length of 128 years. We did not consider autocorrelation, as growth is strongly
167 influenced by hydrological factors, and our goal was to detect long-term effects (Jansma,
168 2020). This processing minimises biological growth effects and retains the climatic signal for
169 spectral analysis (Figure 2, lower panel).

170



171

172

173 Figure 2. Mean ring with and growth index (GI) records of (Angstel/Vecht) NLBOG26 record
174 of Bazelmans et al. (submitted).

175

176 Analyses were performed using functions from the Astrochron (Meyers et al., 2014, 2019)
177 and WaverideR (Arts, 2023; Arts et al., 2024) R packages (see S1 for the R code). The
178 Continuous Wavelet Transform (CWT) and Cross Wavelet Transform (XWT) functions from
179 WaverideR were applied to generate scalograms and examine the temporal variability of
180 cycles. Dominant cycles with high-power spectral peaks were then extracted using the Taner
181 bandpass filter (Taner et al., 2000), as implemented in the Astrochron R package.

182

183 Potential nonlinear interactions among higher-order cycles were evaluated using a modified
184 bicoherence function from the Astrochron R package, which employs WOSA-based
185 smoothing to reduce spurious peaks (Choudhury et al., 2008; Hagelberg et al., 1999; Sullivan
186 et al., 2023; Welch et al., 1967). Bandwidth smoothing was applied to obtain reliable
187 estimates, thereby reducing frequency resolution relative to the original power spectrum.
188 Nonlinear components were subsequently extracted from the GI record using a Taner filter
189 to identify combination tones generated by interactions among higher-order cycles.

190

191 Finally, to assess whether the detected cycles reflect large-scale North Atlantic forcing, a
192 Cross Wavelet Transform (XWT) was computed between the growth index (GI) record
193 NLBOG26 and the Uamh an Tartair stalagmite growth-rate record from northwestern
194 Scotland (Baker et al., 2015), which is known to capture variability related to the Atlantic
195 Multidecadal Oscillation (AMO) and the North Atlantic Oscillation (NAO) (Baker et al., 2015).
196 This step enables the identification of phase relationships and the evaluation of the regional
197 hydroclimatic significance of the detected cyclicity.

198

199 4. Results

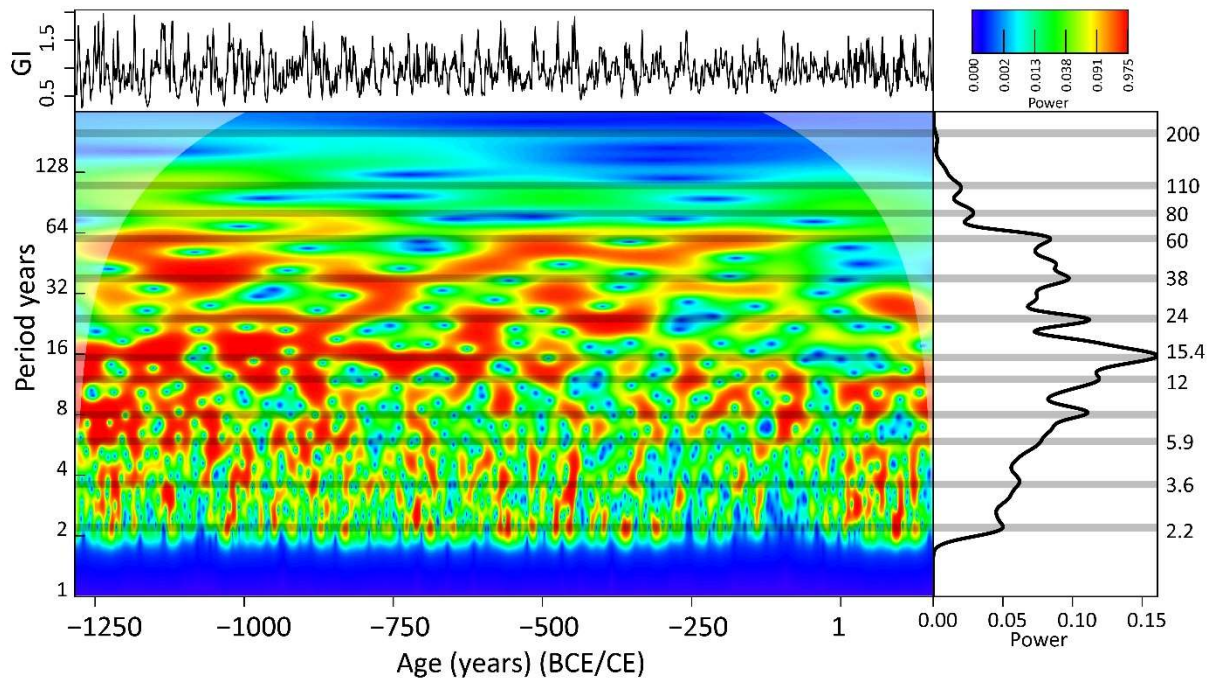
200

201 4.1 Spectral analysis and cycle extraction

202

203 The wavelet scalogram of the Angstel/Vecht chronology reveals spectral peaks with periods
204 of 2.2, 3.6, 5.9, 8, 12, 15.4, 24, 38, 60, 80, 110, and 200 years (Figure 3). This broad spectrum
205 indicates that the Growth Index (GI) was paced by climatic cycles ranging from biannual to
206 centennial cycles.

207



208
209

210 Figure 3. CWT scalogram of the growth index (GI) records of (Angstel/Vecht) NLBOG26
211 record of Bazelmans et al. (submitted).

212

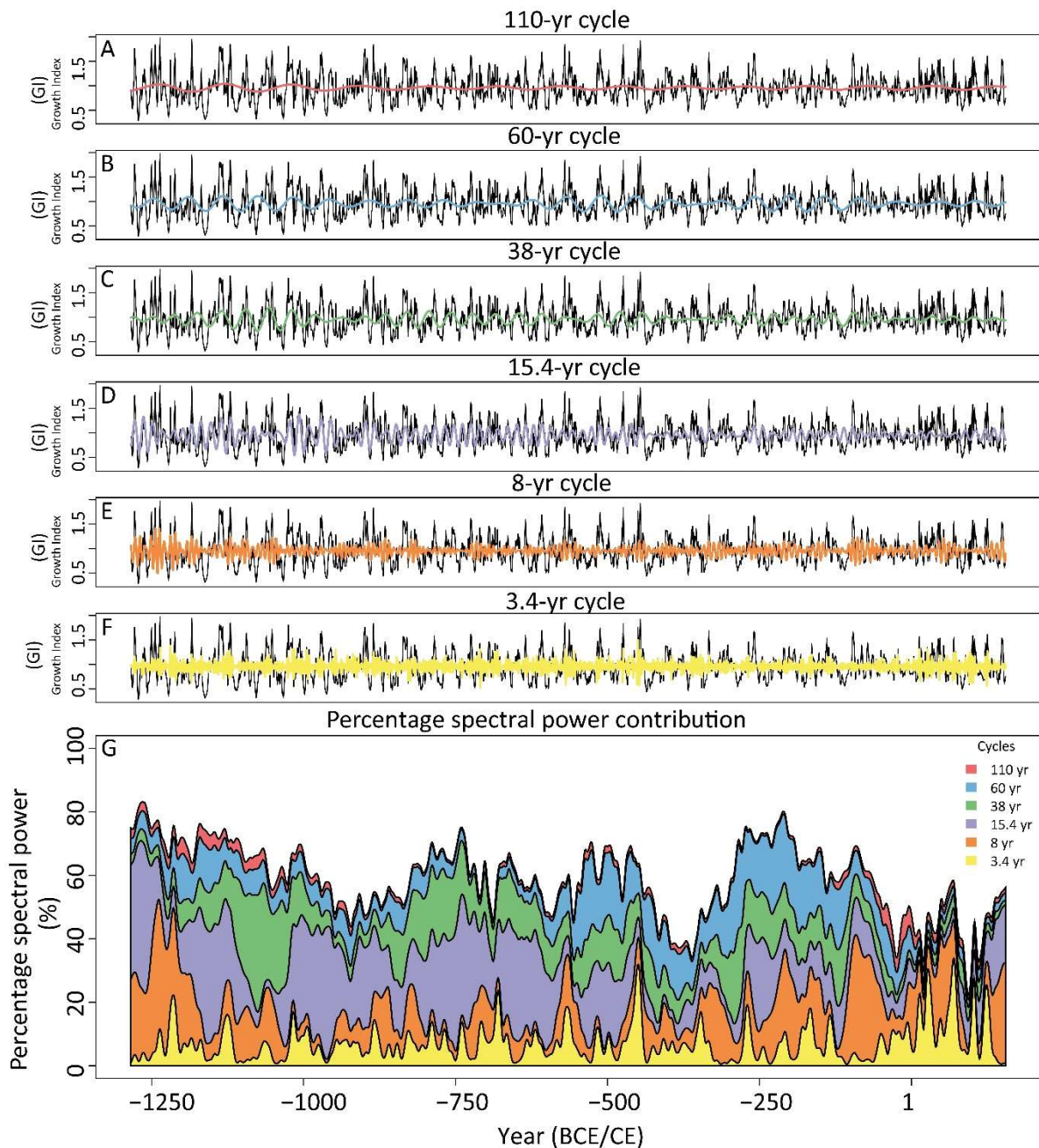
213 To isolate the astronomical signal from the underlying noise, we performed targeted
214 extraction of the most prominent spectral peaks (approximately 3.6, 8, 15, 38, 60, and 110
215 years) from the GI record (Figures 4A-4F). Lower-amplitude spectral peaks, specifically those
216 at 2.2, 5.9, 12, 24, 80, and 200 years, were excluded from further analysis due to insufficient
217 spectral power and a lack of temporal persistence, especially for the shorter-period cycles.

218 The spectral power of each selected cycle was calculated as a percentage of the total
219 spectral power, quantifying its relative contribution over time (Figure 4G). When plotted
220 alongside the original GI record, the 15.4, 38, and 60-year cycles are clearly evident in the
221 raw data. Conversely, the shorter (3.4 and 8-year) and longest (110-year) components are
222 less visually apparent but still contribute to the overall spectral power.

223

224 The relative contribution of each cycle to the total spectral power varies considerably
225 through time (Figure 4G). The ~110-year cycle shows low spectral power, suggesting a minor
226 influence on overall variability. In contrast, the ~60-year cycle dominates the low-frequency
227 domain, accounting for a substantial portion of the spectral power prior to year 1. Cycles of

228 decadal duration (38 and 15.4 years) exhibit distinct yet intermittent contributions,
229 reflecting periodic but non-continuous modulation. The shorter cycles (~8 and ~3.4 years)
230 exhibit low power and high variability, suggesting they represent high-frequency
231 fluctuations superimposed on longer-term trends. Ultimately, the stacked spectral power
232 distribution reveals a hierarchical organisation of cyclicities within the GI record,
233 characterised by temporal shifts in the dominance of different frequency components.
234



235
236 Figure 4. Extracted cycles and spectral power. A GI record and extracted 110-year cycle, B GI
237 record and extracted 60-year cycle, C GI record and extracted 38-year cycle, D GI record and

238 extracted 15-year cycle, E GI record and extracted 8-year cycle, F GI record and extracted
239 3.4-year cycle, G extracted the spectral power (CWT) of the different cycles.

240

241 4.2 Identifying non-linear interactions via bicoherence analysis

242

243 Bicoherence spectral analysis reveals that several shorter-period cycles in the Angstel/Vecht
244 chronology emerge as nonlinear combination tones of longer-period oscillations (Figure 5).

245 Due to the complexity of the bicoherence plot, we have selected the most prominent
246 spectral interactions as representative examples of this nonlinear coupling. Specifically, the
247 interactions between the 110- and 60-year cycles, which generate a ~40-year combination
248 cycle, and the 24- and 40-year cycles, which combine into a ~15-year cycle, are presented as
249 primary examples.

250

251 These specific combinations were selected because they exhibit higher spectral power, and
252 the 45-degree dashed line in the figure can clearly be linked to a spectral peak.

253 Furthermore, these two cases result in cycles closely analogous to the 15.4- and 38-year
254 cycles observed in the CWT scalogram (see Figure 3). While other non-linear combinations
255 are observed, they are poorer examples due to their lower power values and less distinct
256 alignments with the primary spectral peaks in the CWT.

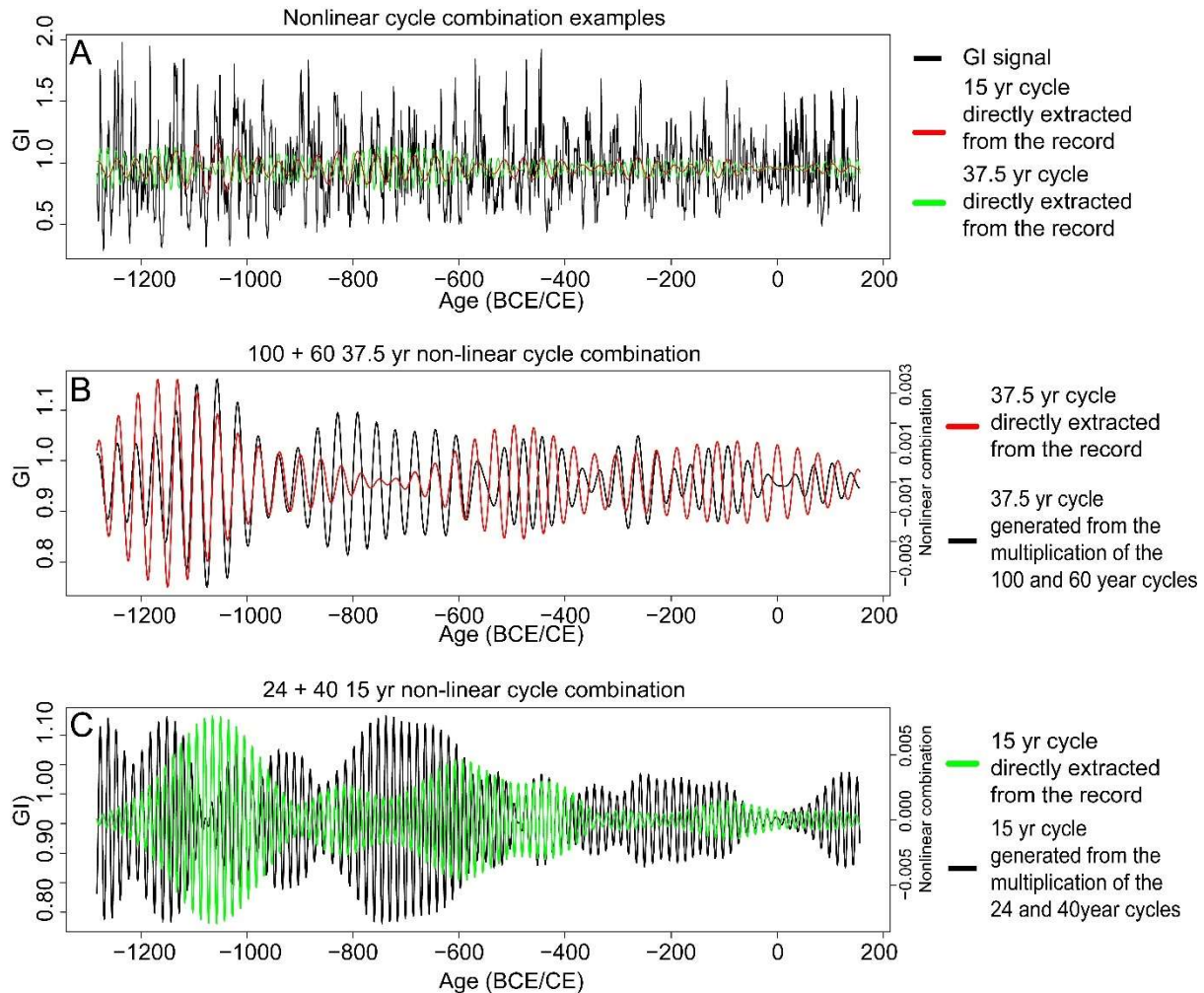
257

258 To evaluate the interpretations drawn from the bicoherence spectra, we synthesised the
259 daughter components by multiplying the two corresponding higher-order parent cycles to
260 determine if they replicate the observed signal (Figure 6). Specifically, the 60- and 110-year
261 combination-tone cycles were multiplied to generate a 37.5-year daughter cycle, which was
262 subsequently extracted from the record (Figure 6B). A similar procedure was applied to the
263 24 and 40-year combination tones to isolate the 15-year daughter cycle (Figure 6C).

264 Comparing the generated daughter cycles with those directly extracted from the raw record
265 shows that they are either completely in phase or completely out of phase. This confirms
266 that nonlinear interactions between cycles created new cyclic components that are either
267 completely in phase or out of phase, depending on the complex nonlinear interaction that
268 occurred during the creation of the daughter cycle. The observation of nonlinear behaviour
269 that creates new cyclic components indicates that tree-ring growth variability is shaped by

282 To better understand the climatic origins of the cycles in the Angstel/Vecht record, a cross-
 283 wavelet analysis was performed against the Uamh an Tartair stalagmite record from
 284 Scotland, an archive known to be driven primarily by North Atlantic weather and its
 285 resulting precipitation patterns (Baker et al., 2011; Baker et al., 2015; Proctor et al., 2000.

286



287

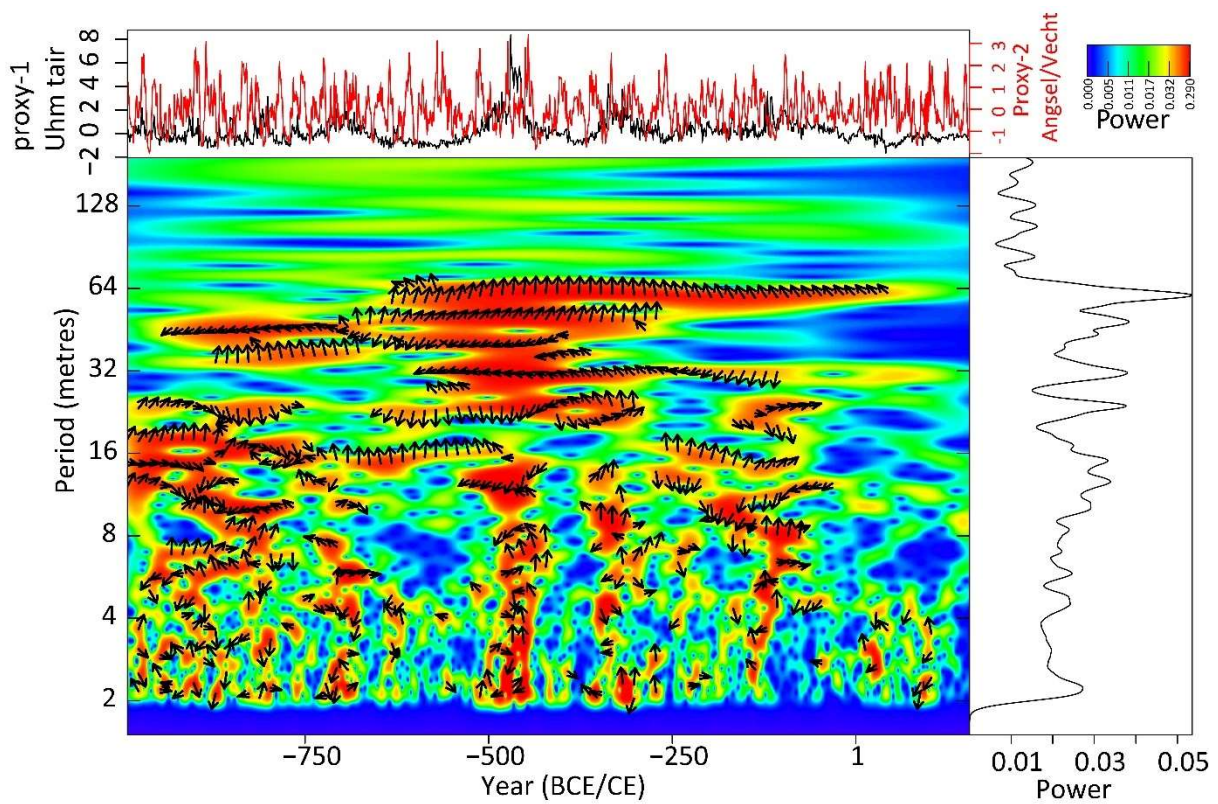
288

289 Figure 6. A. The growth index and the 15- and 37.5-yr cycles directly extracted from the
 290 record, B. 37.5-yr cycle directly extracted from the record, and the 37.5-yr cycle generated
 291 from the 60 and 100-yr cycles, C. 15-yr cycle directly extracted from the record and the 15-
 292 yr cycle generated from the 24 and 40-yr cycles.

293

294 Cross-wavelet analysis between the stalagmite and bog oak GI records identifies significant
 295 shared spectral power across multidecadal to centennial scales. The highest spectral density
 296 is concentrated within the 30-60-year periodicities (Figure 7). A Secondary high-power area

297 is observed between the 30- and 8-year period intervals. In this interval, spectral power is
 298 more variable, as is the direction of the phase arrows. Phase arrows in the cross-wavelet
 299 plots predominantly point upward ($\approx 90^\circ$) or downward ($\approx 270^\circ$), with more upward-pointing
 300 arrows, indicating a persistent quarter-cycle phase-lag. The phase arrows more or less point
 301 either upwards or downwards, regardless of the observed period range, demonstrating that
 302 the stalagmite signal consistently leads the bog oak response by a quarter cycle, regardless
 303 of cycle length. Consequently, the absolute temporal lag scales with the associated
 304 periodicity.
 305



306
 307 Figure 7. cross-wavelet spectra between the Angsel/Vecht bog oak GI dataset and the
 308 stalagmite growth rate dataset from the Uamh an Tartair stalagmite record, Scotland, of
 309 Baker et al., 2015.

310

311 5. Discussion

312

313 5.1 Hydroclimatic Forcing and System Response Mechanisms

314

315 Based on our spectral analyses and cross-wavelet comparisons, we can argue that North
316 Atlantic climate variability exerted a persistent influence on hydroclimate and oak growth
317 across decadal to centennial timescales. The prevalent $\sim 90^\circ$ phase lag observed between
318 the Uamh an Tartair stalagmites and the Angstel/Vecht bog oaks does not reflect a
319 chronological mismatch. Instead, it documents fundamentally different environmental
320 response mechanisms to a shared North Atlantic climatic forcing (Bond et al., 2001; Hurrell,
321 1995; Olsen et al., 2012; Rust et al., 2020; Schultz et al., 2015).

322

323 The Uamh an Tartair stalagmite growth rate primarily reflects changes in precipitation and
324 groundwater recharge in a blanket peat-dominated karst setting (Baker et al., 2015; Proctor
325 et al., 2000). Stalagmite growth is controlled by hydrological flux through the soil-peat
326 column, which governs soil CO₂ production, carbonate dissolution, and drip-water
327 supersaturation (Baker et al., 2011; Baker et al., 2015; Proctor et al., 2000). Because the
328 vadose zone above the cave is only several meters thick and the catchment area is small,
329 precipitation changes are transmitted rapidly into the cave system. Consequently,
330 stalagmite growth rates respond quasi-instantaneously to shifts in moisture balance,
331 particularly during the onset of wetter conditions associated with persistent positive NAO
332 phases (Baker et al., 2011; Baker et al., 2015; Proctor et al., 2000). The stalagmite archive,
333 therefore, records the forcing or instantaneous-flux component of hydroclimatic variability.

334

335 Conversely, the GI record from the Angstel/Vecht delta partly reflects local groundwater
336 conditions governed by the long-term storage and residence times of water within the vast
337 Rhine drainage basin. Sustained increases in precipitation and meltwater lead to a
338 cumulative rise in groundwater levels across the drainage basin. Because the groundwater
339 response is not instantaneous, it represents an integrated signal of basin-wide processes.
340 This includes not only local rainfall but also seasonal snow and glacial melt from the upper
341 reaches of the catchment (Viviroli et al., 2003). As long as the recharge rate remains
342 positive, groundwater levels continue to rise. Therefore, the lowest GI results not from the
343 peak rainfall rate, but from the cumulative multi-year effect that produces the absolute
344 highest groundwater levels. These high levels suppress oak growth by reducing oxygen
345 availability in the root zone, inducing metabolic stress, and inhibiting nutrient uptake (Rust
346 et al., 2020; Sass-Klaassen and Hanraets, 2006; Vreugdenhil et al., 2006).

347

348 This interpretation requires nuance because high-frequency, extreme precipitation and
349 flood events operate on fundamentally different scales than groundwater. Unlike the slow,
350 integrated rise of the water table, these events occur on seasonal or weekly timescales and
351 can trigger near-instantaneous physiological stress in oak growth. These extremes should
352 not be dismissed as stochastic noise; rather, they exhibit distinct periodicities driven by
353 specific phases of the NAO or AMO. Specifically, positive NAO phases and warm AMO states
354 are associated with increased frequency and magnitude of winter floods in the Rhine
355 catchment (Toonen et al., 2016; Scaife et al., 2008). Consequently, these periodic shifts in
356 flood frequency can superimpose a sharp, instantaneous climatic signal onto the broader,
357 long-term integrated groundwater trend.

358

359 In dynamical terms, the tree-ring record represents a state-sensitive response to the
360 integral of the forcing variable. Peak groundwater-induced stress is reached only after a
361 prolonged period of moisture surplus. This integrating behaviour, combined with
362 interference from periodic acute floods, explains the persistent but imperfect quarter-cycle
363 lag observed in the Uamh an Tartair stalagmite growth rate record. While the Uamh an
364 Tartair stalagmite growth reacts to climatic change near-instantaneously, the Dutch bog
365 oaks reflect the cumulative effects of that same change. Given that the Uamh an Tartair
366 stalagmite securely records North Atlantic hydroclimate, we can conclude that the
367 Angstel/Vecht bog oaks are driven by the exact same forcing, albeit expressed through a
368 different hydrological filter. The lowland delta environment functions as a non-linear
369 biological integrator, where external climatic forcings are transformed, amplified, or phase-
370 shifted as they cascade through ocean-atmosphere feedbacks before being filtered through
371 the immense hydrology of the Rhine drainage basin (Bond et al., 2001; Hurrell et al., 2003;
372 Olsen et al., 2012; Scaife et al., 2013).

373

374 5.2 Linking Cycles to Climatic Patterns

375

376 The broad spectrum of periodicities observed in the Angstel/Vecht GI record indicates that
377 lowland oak growth is governed by a hierarchy of interacting climatic patterns, ranging from
378 fast-acting atmospheric oscillations to slow-moving oceanic and solar drivers.

379

380 The shortest spectral peaks in the GI record can be coupled to high-frequency atmospheric
381 variability. Periodicities between 2 and 4 years in the Angstel/Vecht record are consistent
382 with the Quasi-Biennial Oscillation (QBO), which modulates stratospheric wind fields and
383 subsequently affects mid-latitude storminess and precipitation (Baldwin et al., 2001;
384 Brönnimann et al., 2016; Oz et al., 2009; Sonett et al., 1992). Cycles between 5.9 and 8 years
385 fall within the characteristic range of the North Atlantic Oscillation (NAO) (Adolph et al.,
386 2024; Appenzeller et al., 1998; Cook et al., 1998; Hurrell, 1995; Luterbacher et al., 2001;
387 Rust et al., 2020, 2022; Wanner et al., 2001). Positive NAO phases enhance westerly airflow,
388 significantly increasing precipitation across northwestern Europe (Knudsen et al., 2011;
389 Sutton and Hodson, 2005). In the Rhine catchment, these NAO- and QBO-driven
390 periodicities dictate the frequency of instantaneous hydrological shocks and extreme winter
391 flood events, which acutely constrain local oak growth (Sass-Klaassen and Hanraets, 2006).

392

393 Underlying these high-frequency atmospheric variations are slower, multidecadal-to-
394 centennial modes of 38, 60 and 110 years in the Angstel/Vecht record that reflect internal
395 variability of the Atlantic Multidecadal Oscillation (AMO) (Ait Brahim et al., 2018; Knudsen et
396 al., 2011). The AMO exhibits a characteristic 60-80-year cyclicity with minor components at
397 35, 85-100, and 110 years (Delworth and Mann, 2000; Enfield et al., 2001; Goldenberg et al.,
398 2001; Gray et al., 2004; Hubeny et al., 2006; Schlesinger and Ramankutty, 1994; Sutton and
399 Hodson, 2005). By modulating North Atlantic sea surface temperatures, the AMO exerts a
400 profound control over continental evaporation and sustained moisture transport into the
401 European interior. Consequently, AMO dynamics drive long-term cumulative groundwater
402 trends that govern the integrated hydrological baseline of the deltaic environment (Baker et
403 al., 2011; Geirsdóttir et al., 2016; Sass-Klaassen and Hanraets, 2006).

404

405 These atmospheric and oceanic systems are further paced by exogenous solar forcing.
406 Periodicities at 12, 24, 58, 80, and 110 years in the Angstel Vecht record can be linked to
407 known solar activity cycles and their amplitude-modulating components, including the 11-
408 year Schwabe and 22-year Hale cycles (Brehm et al., 2021; Hathaway, 2015; Land et al.,
409 2020; Miyahara et al., 2009; Murphy et al., 1994; Nintcheu-Fata et al., 2003; Sonett and
410 Finney, 1990). Imperfect period alignment of the 12- and 24-year cycles with solar cycles

411 either reflects a deep-time period shift or reflects the thermal inertia of the oceans and
412 delayed propagation through the atmosphere, which alters the period of the solar signal
413 (Scaife et al., 2013; Seip et al., 2019). Solar forcing shapes regional hydroclimate via
414 atmospheric circulation patterns, including stratosphere-troposphere coupling, which
415 affects the stability of NAO-like pressure patterns (Hathaway and Wilson, 2004; Lohmann et
416 al., 2004; Seidenglanz et al., 2012; Seip et al., 2019).

417

418 While individual cycles can be linked to discrete climatic phenomena, our bicoherence
419 analyses demonstrate that regional hydroclimate arises from nonlinear interactions among
420 climatic cycles (Land et al., 2020; Schultz et al., 2015). Combination tones, such as the 38-
421 year cycle (arising from the non-linear interaction between the 100- and 60-year
422 interactions) and the 15-year cycle (arising from the non-linear interaction between the 24-
423 and 40-year cycles), illustrate that these forcings do not act in isolation. Instead, slow-
424 moving oceanic boundary conditions (AMO) and top-down solar modulation interact to
425 dictate the probability and intensity of high-frequency atmospheric climatic cycles (NAO and
426 QBO).

427

428 Consequently, the spectral architecture of the Angstel/Vecht chronology is not a simple
429 additive record; it captures interacting climatic modes in which multidecadal oceanic and
430 solar trends systematically modulate high-frequency atmospheric dynamics that govern
431 extreme flood events and groundwater recharge in northwestern Europe. This complex
432 interaction explains why the observed spectral peaks in the Angstel/Vecht record
433 sometimes exhibit slight offsets from the pure solar or oceanic cycles that drive them. The
434 clear expression of both individual cycles and their nonlinear interactions is closely tied to
435 the localised sampling strategy of the Angstel/Vecht chronology, which targets a single
436 drainage region and groundwater system within one branch of the Rhine. By minimising
437 spatial averaging, this approach reduced signal attenuation that might have obscured
438 cyclical variability in larger regional composites. At the same time, it highlights a key
439 limitation. Although a shared North Atlantic forcing links the Angstel/Vecht oak record to
440 other archives, such as the Uamh an Tartair stalagmites, site-specific hydrological filtering
441 prevents straightforward one-to-one correlation of individual cycles between records.
442 Hydroclimatic responses, therefore, remain inherently site-sensitive. This raises an

443 important question: whether the identified nonlinear interactions are broadly transferable
444 across the region or instead reflect local characteristics of the Angstel/Vecht region.
445 Addressing this will require systematic comparisons across contrasting drainage systems,
446 which are essential for determining whether such nonlinear coupling is a fundamental
447 feature of North Atlantic climate dynamics and for understanding how these interactions
448 would have propagated through and are recorded in palaeoclimatological records within the
449 wider Rhine drainage basin.

450

451 6. Conclusion

452

453 The nearly 1.5-millennia-long Angstel/Vecht bog oak chronology provides a highly sensitive
454 archive of hydroclimatic variability in northwestern Europe, as evidenced by wavelet, cross-
455 wavelet, and bicoherence spectral analyses, as well as Taner filtering. We show that tree
456 growth in this lowland wetland system is governed by a broad spectrum of cyclicities
457 ranging from interannual to centennial timescales, reflecting the combined influence of
458 atmospheric, oceanic, and solar forcing mechanisms.

459

460 The identified periodicities closely match known modes of North Atlantic climate variability,
461 including the North Atlantic Oscillation and the Atlantic Multidecadal Oscillation, as well as
462 solar cycles. However, the spectral structure of the tree-ring record is not a simple linear
463 imprint of these external forcings. Instead, bicoherence analysis reveals that a significant
464 portion of the observed variability arises from nonlinear interactions among primary cycles,
465 generating combination tones that contribute substantially to the overall signal. This
466 highlights the importance of considering nonlinear dynamics when interpreting
467 dendrochronological records.

468

469 The cross-wavelet comparison with the Uamh an Tartair stalagmite record provides
470 compelling evidence that both archives are driven by a common North Atlantic
471 hydroclimatic forcing. The persistent quarter-cycle phase lag between the records reflects
472 contrasting system responses rather than chronological discrepancies. While the stalagmite
473 responds rapidly to changes in precipitation, the Angstel/Vecht oak record integrates
474 hydroclimatic forcing over longer timescales through groundwater storage and catchment-

475 scale hydrological processes. This integrative behaviour results in a quarter-cycle delayed
476 response, demonstrating that lowland wetland trees act as nonlinear biological filters of
477 climate variability.

478

479 Our findings underscore that hydroclimate in the Rhine–Meuse delta is shaped by a
480 hierarchy of interacting processes, in which high-frequency atmospheric variability operates
481 within the boundary conditions set by slower oceanic and solar cycles. The Angstel/Vecht
482 record captures both the immediate impacts of extreme hydrological events and the
483 cumulative effects of sustained moisture anomalies, offering a unique perspective on how
484 climate signals are transformed within terrestrial systems. Overall, this work highlights the
485 untapped potential of lowland bog oak archives to reconstruct past hydroclimatic variability
486 and to disentangle the complex interplay of forcing mechanisms across timescales.

487

488 7. Data availability

489 No new data was generated as part of this study.

490

491 8. Supplementary material

492 S1: The R code used in this study can be found at [https://github.com/stratigraphy/Weesp-](https://github.com/stratigraphy/Weesp-dendro-cycles)
493 [dendro-cycles](https://github.com/stratigraphy/Weesp-dendro-cycles).

494

495 9. References

496 Adolph, M. L., Czerwiński, S., Dreßler, M., Strobel, P., Bliedtner, M., Lorenz, S., Debret, M.,
497 and Haberzettl, T., 2024. North Atlantic Oscillation polarity during the past 3000 years
498 derived from sediments of a large lowland lake, Schweriner See, in NE Germany. *Climate of*
499 *the Past*, 20, 2143-2165. <https://doi.org/10.5194/cp-20-2143-2024>

500 Ait Brahim, Y., Wassenburg, J. A., Cruz, F. W., Sifeddine, A., Scholz, D., Bouchaou, L., Dassié,
501 E. P., Jochum, K. P., Edwards, R. L., and Cheng, H., 2018. Multi-decadal to centennial hydro-
502 climate variability and linkage to solar forcing in the Western Mediterranean during the last
503 1000 years. *Scientific Reports*, 8, 17446. <https://doi.org/10.1038/s41598-018-35498-x>

504 Appenzeller, C., Stocker, T. F., and Anklin, M., 1998. North Atlantic Oscillation dynamics
505 recorded in Greenland ice cores. *Science*, 282(5388), 446-449.

506 Arts, M., 2023. WaverideR: Extracting Signals from Wavelet Spectra. R package version
507 0.1.0. Available on CRAN: <https://CRAN.R-project.org/package=WaverideR>

508 Arts, M., Corradini, C., Pondrelli, M., Pas, D., and Da Silva, A. C., 2024. Age and orbital forcing
509 in the upper Silurian Cellon section (Carnic Alps, Austria) uncovered using the WaverideR R
510 package. *Frontiers in Earth Science*, 12, 1357751.
511 <https://doi.org/10.3389/feart.2024.1357751>

512 Baker, A., Hellstrom, J., Kelly, B. F. J., Mariethoz, G., and Trouet, V., 2015. A composite
513 annual-resolution stalagmite record of North Atlantic climate over the last three millennia.
514 *Scientific Reports*, 5, 10307. <https://doi.org/10.1038/srep10307>

515 Baker, A., Wilson, R., Fairchild, I. J., Franke, J., Spötl, C., Matthey, D., Trouet, V., and Fuller, L.,
516 2011. High resolution $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$ records from an annually laminated Scottish
517 stalagmite and relationship with last millennium climate. *Global and Planetary Change*, 79,
518 303-311. <https://doi.org/10.1016/j.gloplacha.2010.12.007>

519 Baldini, L. M., McDermott, F., Foley, A. M., and Baldini, J. U. L., 2008. Spatial variability in the
520 European winter precipitation $\delta^{18}\text{O}$ -NAO relationship: Implications for reconstructing NAO-
521 mode climate variability in the Holocene. *Geophysical Research Letters*, 35, L04709.
522 <https://doi.org/10.1029/2007GL032027>

523 Baldwin, M. P., et al., 2001. The quasi-biennial oscillation. *Reviews of Geophysics*, 39(2),
524 179-229.

525 Bond, G., Kromer, B., Beer, J., Muscheler, R., Evans, M. N., Showers, W., Hoffmann, S., Lotti-
526 Bond, R., Hajdas, I., and Bonani, G., 2001. Persistent solar influence on North Atlantic
527 climate during the Holocene. *Science*, 294(5549), 2130-2136.

528 Börgel, F., Frauen, C., Neumann, T., and Markus Meier, H. E., 2020. The atlantic multidecadal
529 oscillation controls the impact of the north atlantic oscillation on north european climate.
530 *Environmental Research Letters*, 15, 104025. <https://doi.org/10.1088/1748-9326/aba925>

531 Bos, I. J., Feiken, H., Bunnik, F., and Schokker, J., 2009. Influence of organics and clastic lake
532 fills on distributary channel processes in the distal Rhine-Meuse delta (The Netherlands).
533 *Palaeogeography, Palaeoclimatology, Palaeoecology*, 284(3-4), 355-374.

534 Brehm, N., et al., 2021. Eleven-year solar cycles over the last millennium revealed by
535 radiocarbon in tree rings. *Nature Geoscience*, 14(1), 10-15.

536 Brönnimann, S., Franke, J., Valler, V., Hand, R., Samakinwa, E., Lundstad, E., Burgdorf, A. M.,
537 Lipfert, L., Pfister, L., Imfeld, N., and Rohrer, M., 2025. Past hydroclimate extremes in Europe

538 driven by Atlantic jet stream and recurrent weather patterns. *Nature Geoscience*, 18(3),
539 246-253.

540 Brönnimann, S., Malik, A., Stickler, A., Wegmann, M., Raible, C. C., Muthers, S., Anet, J. G.,
541 Rozanov, E. V., and Schmutz, W. K., 2016. Multidecadal variations of the effects of the
542 Quasi-Biennial Oscillation on the climate system. *Atmospheric Chemistry and Physics*, 16,
543 15529-15543.

544 Brouwer, F., van Delft, S. P. J., and Kemmers, R. H., 2002. Landinventarisatie en ruimtelijke
545 systeemanalyse van het herinrichtingsgebied De Vechtstreek, fase 2; resultaten van een
546 bodemgeografisch onderzoek.

547 Choudhury, S. M., Shah, S. L., and Thornhill, N. F., 2008. Bispectrum and bicoherence. In
548 *Diagnosis of Process Nonlinearities and Valve Stiction: Data Driven Approaches*, Springer,
549 29-41.

550 Coder, K.D., 1994. Flood damage to trees. University of Georgia Extension, Forest Resources
551 Publication FOR94-61.

552 Cook, E. R., 1990. A conceptual linear aggregate model for tree rings. In Cook, E. R., and
553 Kairiukstis, L. A. (Eds.), *Methods of Dendrochronology*. Kluwer Academic Publishers,
554 Dordrecht, 98-104.

555 Cook, E. R., D'Arrigo, R. D., and Briffa, K. R., 1998. A reconstruction of the North Atlantic
556 Oscillation using tree-ring chronologies from North America and Europe. *The Holocene*, 8(1),
557 9-17.

558 Copini, P., den Ouden, J., Robert, E.M.R., Tardif, J.C., Loesberg, W.A., Goudzwaard, L., Sass-
559 Klaassen, U., 2016. Flood-ring formation and root development in response to experimental
560 flooding of young *Quercus robur* trees. **Front. Plant Sci.** 7, 775, 1–14.
561 <https://doi.org/10.3389/fpls.2016.00775>.

562 De Lange, G. W., and Bles, B., 1963. De bodemgesteldheid rondom de gemeente Weesp:
563 structuurplan Weesp.

564 Delworth, T. L., and Mann, M. E., 2000. Observed and simulated multidecadal variability in
565 the Northern Hemisphere. *Climate Dynamics*, 16, 661-676.

566 Denniston, R. F., et al., 2018. A stalagmite test of North Atlantic SST and Iberian
567 hydroclimate linkages over the last two glacial cycles. *Climate of the Past*, 14(12), 1893-
568 1913.

569 Dieppois, B., Durand, A., Fournier, M., and Massei, N., 2013. Links between multidecadal
570 and interdecadal climatic oscillations in the North Atlantic and regional climate variability of
571 northern France and England since the 17th century. *Journal of Geophysical Research:*
572 *Atmospheres*, 118, 4359-4372. <https://doi.org/10.1002/jgrd.50392>

573 Domeisen, D., Badin, G., and Koszalka, I., 2018. How Predictable Are the Arctic and North
574 Atlantic Oscillations? Exploring the Variability and Predictability of the Northern
575 Hemisphere. *Journal of Climate*, 31, 997-1014.

576 Eade, R., Stephenson, D. B., Scaife, A. A., and Smith, D. M., 2022. Quantifying the rarity of
577 extreme multi-decadal trends: How unusual was the late twentieth century trend in the
578 North Atlantic Oscillation? *Climate Dynamics*, 58, 1555-1568.

579 Enfield, D. B., Mestas-Nunez, A. M., and Trimble, P. J., 2001. The Atlantic multidecadal
580 oscillation and its relation to rainfall and river flows in the continental US. *Geophysical*
581 *Research Letters*, 28, 2077-2080.

582 Faust, J. C., Fabian, K., Milzer, G., Giraudeau, J., and Knies, J., 2016. Norwegian fjord
583 sediments reveal NAO related winter temperature and precipitation changes of the past
584 2800 years. *Earth and Planetary Science Letters*, 435, 84-93.
585 <https://doi.org/10.1016/j.epsl.2015.12.003>

586 Folland, C. K., Knight, J., Linderholm, H. W., Fereday, D., Ineson, S., and Hurrell, J. W., 2009.
587 The summer North Atlantic Oscillation: Past, present, and future. *Journal of Climate*, 22,
588 1082-1103.

589 Goldenberg, S. B., Landsea, C. W., Mestas-Nuñez, A. M., and Gray, W. M., 2001. The recent
590 increase in Atlantic hurricane activity: Causes and implications. *Science*, 293, 474-479.

591 Geirsdóttir, Á., Miller, G.H., Larsen, D.J. and Ólafsdóttir, S., 2013. Abrupt Holocene climate
592 transitions in the northern North Atlantic region recorded by synchronized lacustrine
593 records in Iceland. *Quaternary Science Reviews*, 70, 48-62.

594 Gray, S. T., Graumlich, L. J., Betancourt, J. L., and Pederson, G. T., 2004. A tree-ring based
595 reconstruction of the Atlantic Multidecadal Oscillation since 1567 AD. *Geophysical Research*
596 *Letters*, 31, L12205.

597 Hagelberg, T., Piasias, N., and Elgar, S., 1991. Linear and nonlinear couplings between orbital
598 forcing and the marine $\delta^{18}\text{O}$ record during the Late Neocene. *Paleoceanography*, 6, 729-746.

599 Hathaway, D. H., 2015. The Solar Cycle. *Living Reviews in Solar Physics*, 12, 4.
600 <https://doi.org/10.1007/lrsp-2015-4>

601 Hathaway, D. H., and Wilson, R. M., 2004. What the Sunspot Record Tells Us About Space
602 Climate. *Solar Physics*, 224, 5-19.

603 Holman, I.P., Rivas-Casado, M., Howden, N.J.K., Bloomfield, J.P. and Williams, A.T. (2009),
604 Linking North Atlantic ocean–atmosphere teleconnection patterns and hydrogeological
605 responses in temperate groundwater systems. *Hydrol. Process.*, 23: 3123-3126.
606 <https://doi.org/10.1002/hyp.7466>

607 Hijma, M. P., and Cohen, K. M., 2019. Holocene sea-level database for the Rhine-Meuse
608 Delta, The Netherlands: Implications for the pre-8.2 ka sea-level jump. *Quaternary Science*
609 *Reviews*, 214, 68-86.

610 Hubeny, J. B., King, J. W., and Santos, A., 2006. Subdecadal to multidecadal cycles of Late
611 Holocene North Atlantic climate variability preserved by estuarine fossil pigments. *Geology*,
612 34, 569-572.

613 Hurrell, J. W., 1995. Decadal Trends in the North Atlantic Oscillation: Regional Temperatures
614 and Precipitation. *Science*, 269(5224), 676-679.

615 Hurrell, J. W., Kushnir, Y., Ottersen, G., and Visbeck, M., 2003. An overview of the North
616 Atlantic oscillation. *Geophysical Monograph-American Geophysical Union*, 134, 1-36.

617 Hurrell, J. W., and Van Loon, H., 1997. Decadal variations in climate associated with the
618 North Atlantic Oscillation. *Climatic Change*, 36, 301-326.

619 Jansma, E., 2020. Hydrological disasters in the NW-European Lowlands during the first
620 millennium AD: A dendrochronological reconstruction. *Netherlands Journal of Geosciences*,
621 99, e11.

622 Kelly, P. M., Leuschner, H. H., Briffa, K. R., and Harris, I. C., 2002. The climatic interpretation
623 of pan-European signature years in oak ring-width series. *The Holocene*, 12, 689-694.

624 Knudsen, M. F., Seidenkrantz, M.-S., Jacobsen, B. H., and Kuijpers, A., 2011. Tracking the
625 Atlantic Multidecadal Oscillation through the last 8,000 years. *Nature Communications*, 2,
626 178.

627 Land, A., Kromer, B., Remmele, S., Brehm, N., and Wacker, L., 2020. Complex imprint of
628 solar variability on tree rings. *Environmental Research Communications*, 2, 101003.

629 Leuschner, H. H., Sass-Klaassen and Hanraets, U., Jansma, E., Baillie, M. G. L., and Spurk, M.,
630 2002. Subfossil European bog oaks: Population dynamics and long-term growth depressions
631 as indicators of changes in the Holocene hydro-regime and climate. *The Holocene*, 12, 695-
632 706.

633 Lohmann, G., Rimbu, N., and Dima, M., 2004. Climate signature of solar irradiance
634 variations: Analysis of long-term instrumental, historical, and proxy data. *International*
635 *Journal of Climatology*, 24, 1045-1056.

636 Lüdecke, H., Cina, R., Dammschneider, H., and Lüning, S., 2020. Decadal and multidecadal
637 natural variability in European temperature. *Journal of Atmospheric and Solar-Terrestrial*
638 *Physics*, 205, 105294.

639 Luterbacher, J., et al., 2001. Extending North Atlantic oscillation reconstructions back to
640 1500. *Atmospheric Science Letters*, 2, 114-124.

641 Massei, N., Durand, A., Deloffre, J., Dupont, J.-P., Valdes, D., and Laignel, B., 2007.
642 Investigating possible links between the North Atlantic Oscillation and rainfall variability in
643 northwestern France over the past 35 years. *Journal of Geophysical Research: Atmospheres*,
644 112, 2005JD007000.

645 Meyers, S. R., 2014. astrochron: A Computational Tool for Astrochronology. R package
646 version 1.6. Available on CRAN: <https://CRAN.R-project.org/package=astrochron>

647 Meyers, S. R., 2019. Cyclostratigraphy and the problem of astrochronologic testing. *Earth-*
648 *Science Reviews*, 190, 190-223.

649 Miyahara, H., Yokoyama, Y., and Yamaguchi, Y. T., 2009. Influence of the Schwabe/Hale solar
650 cycles on climate change during the Maunder Minimum. *Proceedings of the International*
651 *Astronomical Union*, 5, 427-433.

652 Murphy, J. O., Sampson, H., Veblen, T. T., and Villalba, R., 1994. Regression Model for the
653 22-year Hale Solar Cycle Derived from High Altitude Tree-ring Data. *Publications of the*
654 *Astronomical Society of Australia*, 11, 157-163.

655 Nintcheu-Fata, S., et al., 2003. Chronomics of tree rings for chronoastrobiology and beyond.
656 *Biomedicine & Pharmacotherapy*, 57, 24-30.

657 Ólafsdóttir, K. B., Geirsdóttir, Á., Miller, G. H., and Larsen, D. J., 2013. Evolution of NAO and
658 AMO strength and cyclicity derived from a 3-ka varve-thickness record from Iceland.
659 *Quaternary Science Reviews*, 69, 142-154.

660 Olsen, J., Anderson, N. J., and Knudsen, M. F., 2012. Variability of the North Atlantic
661 Oscillation over the past 5,200 years. *Nature Geoscience*, 5, 808-812.

662 Oz, A., Sen, A., Sancho, C., and Genty, D., 2009. Wavelet analysis of late holocene stalagmite
663 records from ortigosa caves in northern Spain. *Journal of Cave and Karst Studies*, 71.

664 Pilcher, J. R., Hillam, J., Baillie, M. G. L., and Pearson, G. W., 1977. A long subfossil oak tree-
665 ring chronology from the north of Ireland. *New Phytologist*, 79, 713-729.

666 Pinto, J. G., and Raible, C. C., 2012. Past and recent changes in the North Atlantic oscillation.
667 *WIREs Climate Change*, 3, 79-90.

668 Poelman, J. N. B., 1966. De bodem van Utrecht: toelichting bij blad 6 van de bodemkaart van
669 Nederland schaal 1:200.000. Stichting voor Bodemkartering, Wageningen.

670 Proctor, C. J., Baker, A., and Barnes, W. L., 2002. A three thousand year record of North
671 Atlantic climate. *Climate Dynamics*, 19, 449-454.

672 Proctor, C. J., Baker, A., Barnes, W. L., and Gilmour, M. A., 2000. A thousand year
673 speleothem proxy record of North Atlantic climate from Scotland. *Climate Dynamics*, 16,
674 815-820.

675 Roig, F. A., Barriopedro, D., García Herrera, R., Patón Dominguez, D., and Monge, S., 2009.
676 North atlantic oscillation signatures in western iberian tree-rings. *Geografiska Annaler:*
677 *Series A, Physical Geography*, 91, 141-157.

678 Rossi, A., Massei, N., and Laignel, B., 2011. A synthesis of the time-scale variability of
679 commonly used climate indices using continuous wavelet transform. *Global and Planetary*
680 *Change*, 78, 1-13.

681 Rust, W., Cuthbert, M., Bloomfield, J., Corstanje, R., Howden, N., and Holman, I., 2020.
682 Exploring the role of hydrological pathways in modulating North Atlantic Oscillation (NAO)
683 teleconnection periodicities from UK rainfall to streamflow. *Hydrology and Earth System*
684 *Sciences*, <https://doi.org/10.5194/hess-2020-312>

685 Rust, W., Bloomfield, J. P., Cuthbert, M., Corstanje, R., and Holman, I., 2022. The importance
686 of non-stationary multiannual periodicities in the North Atlantic Oscillation index for
687 forecasting water resource drought. *Hydrology and Earth System Sciences*, 26, 2449-2467.

688 Sass-Klaassen and Hanraets, U., and Hanraets, E., 2006. Woodlands of the past: The
689 excavation of wetland woods at Zwolle-Stadshagen (the Netherlands). Growth pattern and
690 population dynamics of oak and ash. *Netherlands Journal of Geosciences*, 85(1), 61-71.

691 Scaife, A., Folland, C., Alexander, L., Moberg, L.V. Knight, ., 2008. European Climate
692 Extremes and the North Atlantic Oscillation. *Journal of Climate*. 21. 72-83.
693 10.1175/2007JCLI1631.1.

694 Schlesinger, M. E., and Ramankutty, N., 1994. An oscillation in the global climate system of
695 period 65-70 years. *Nature*, 367, 723-726.

696 Schultz, J. A., et al., 2015. Sensitivity of proxies on non-linear interactions in the climate
697 system. *Scientific Reports*, 5, 18560.

698 Seidenglanz, A., Prange, M., Varma, V., and Schulz, M., 2012. Ocean temperature response
699 to idealized Gleissberg and de Vries solar cycles in a comprehensive climate model.
700 *Geophysical Research Letters*, 39, L20702.

701 Seip, K. L., Grøn, Ø., and Wang, H., 2019. The North Atlantic Oscillations: Cycle Times for the
702 NAO, the AMO and the AMOC. *Climate*, 7, 43.

703 Sonett, C. P., and Finney, S. A., 1990. The spectrum of radiocarbon. *Philosophical*
704 *Transactions of the Royal Society of London. Series A*, 330, 413-426.

705 Sonett, C. P., Williams, C. R., and Mörner, N.-A., 1992. The fourier spectrum of Swedish
706 riverine varves: Evidence of sub-arctic quasi-biennial (QBO) oscillations. *Global and*
707 *Planetary Change*, 6, 57-65.

708 Steirou, E., Gerlitz, L., Apel, H., and Merz, B., 2017. Links between large-scale circulation
709 patterns and streamflow in Central Europe: A review. *Journal of Hydrology*, 549, 484-500.

710 Sullivan, N. B., et al., 2023. Millennial-scale variability of the Antarctic ice sheet during the
711 early Miocene. *Proceedings of the National Academy of Sciences (PNAS)*, 120(19),
712 e2215036120.

713 Sutton, R. T., and Hodson, D. L. R., 2005. Atlantic Ocean forcing of North American and
714 European summer climate. *Science*, 309, 115-118.

715 Taner, M. T., 2000. *Attributes revisited*. Technical Publication.

716 Toonen, W. H. J., Middelkoop, H., Konijnendijk, T. Y. M., Macklin, M. G., and Cohen, K. M.
717 (2016) The influence of hydroclimatic variability on flood frequency in the Lower Rhine.
718 *Earth Surf. Process. Landforms*, 41: 1266–1275. doi: 10.1002/esp.3953.

719 Trouet, V., et al., 2009. Persistent positive North Atlantic Oscillation mode during the
720 Medieval Climate Anomaly. *Science*, 324, 78-80.

721 Van Asselen, S., Cohen, K. M., and Stouthamer, E., 2017. The impact of avulsion on
722 groundwater level and peat formation in delta floodbasins during the middle-Holocene
723 transgression in the Rhine-Meuse delta, The Netherlands. *The Holocene*, 27(11), 1694-1706.

724 Van der Woude, J. D., 1984. The fluviolagoonal palaeoenvironment in the Rhine/Meuse
725 deltaic plain. *Sedimentology*, 31, 395-400.

726 Van Dinter, M. V., Cohen, K. M., Hoek, W. Z., et al., 2017. Late Holocene lowland fluvial
727 archives and geoarchaeology: Utrecht's case study of Rhine River abandonment under
728 Roman and medieval settlement. *Quaternary Science Reviews*, 166, 227-265.

729 Viviroli, D., Weingartner, R. and Messerli, B., 2003. Assessing the hydrological significance of
730 the world's mountains. *Mountain research and Development*, 23(1), 32-40.

731 Vos, P., 2015. Origin of the Dutch coastal landscape: long-term landscape evolution of the
732 Netherlands during the Holocene, described and visualized in national, regional and local
733 palaeogeographical map series. PhD Thesis, Utrecht University.

734 Vos, P., van der Meulen, M., Weerts, H., and Bazelmans, J., 2020. *Atlas of the Holocene
735 Netherlands: Landscape and Habitation since the Last Ice Age*. Amsterdam University Press.

736 Vreugdenhil, S. J., Kramer, K., Pelsma, T., 2006. Effects of flooding duration, -frequency and
737 -depth on the presence of saplings of six woody species in north-west Europe, *Forest
738 Ecology Management*, vol. 236, 47-55

739 Wanner, H., et al., 2001. North Atlantic Oscillation-concepts and studies. *Surveys in
740 Geophysics*, 22, 321-381.

741 Weerts, H., Cleveringa, P., and Gouw, M., 2002. De Vecht/Angstel, een riviersysteem in het
742 veen. *Grondboor & Hamer*, 56(3/4), 66-71.

743 Welch, P. D., 1967. The use of fast Fourier transform for the estimation of power spectra: A
744 method based on time averaging over short, modified periodograms. *IEEE Transactions on
745 Audio and Electroacoustics*, 15(2), 70-73.