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1 **A comprehensive study of the CO₂-O₂ isotope exchange technique to standardize triple**
2 **oxygen measurements in CO₂**

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33 **Abstract**

34 Triple oxygen isotope ($\Delta^{17}\text{O}$) measurements in CO_2 using the platinum catalysed $\text{CO}_2\text{-O}_2$
35 exchange method require precise determination of oxygen isotope exchange fractionation
36 factors ($^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$). These catalysed exchange fractionation factors are typically
37 determined at elevated temperature (~ 750 °C). However, large inter-laboratory inconsistencies
38 have been reported in these factors, primarily attributed to thermal-gradient-induced
39 fractionation arising from portions of the exchange reactor extending beyond the heated zone
40 (hereafter referred to as the cold zone). This study experimentally determines
41 $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$, $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and triple oxygen isotope exponent ($\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$) by systematically reducing
42 cold zone over a temperature range of 600 to 1000 °C using novel setup alongside conventional
43 exchange reactor configurations. The novel setup enabled uniform heating of $\text{CO}_2\text{-O}_2$ mixture,
44 effectively eliminating thermal gradient and yielded exchange fractionation factors and
45 exponent in near-perfect agreement with the revised theoretical equilibrium values. In the
46 conventional reactors, reduction of the cold zone improved agreement with equilibrium values;
47 nevertheless, notable discrepancies from theoretical predictions remained even at cold zone
48 proportion of 25%. A crossover between $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ ($^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}} > ^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$) was
49 observed beyond certain temperature depending on the cold zone, reflecting kinetic effects
50 induced by thermal gradient. These results establish that the theory-experimental discrepancies
51 in exchange fractionation factors are driven by thermal gradient and reveal a robust exponential
52 correlation between these factors and hot zone volumes. Here, we propose a new framework
53 to standardize fractionation factors accounting for the hot and cold zones of the exchange
54 reactor, which is crucial for inter-laboratory consistent $\Delta^{17}\text{O}$ measurements in CO_2 .

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56 **Keywords:** Triple oxygen isotope, fractionation factor, thermal diffusion, cold zone

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63 1. Introduction

64 The deviations from the expected proportionality between $\delta^{18}\text{O}$ and $\delta^{17}\text{O}$, though small, have
65 the potential to provide deeper insights into the reaction pathways involved in both mass-
66 dependent and mass-independent realms (Thiemens and Heidenreich, 1983; Thiemens et al.,
67 2012; Bao et al., 2016; Miller and Pack, 2021a). Focusing on the ^{17}O anomaly in CO_2 , the
68 signature is increasingly being utilised to estimate stratosphere-troposphere exchange fluxes,
69 atmospheric CO_2 budget, terrestrial and marine gross primary productivity (Thiemens et al.,
70 2014; Laskar et al., 2016, 2019, 2020; Liang et al., 2017a, 2017b, 2023; Koren et al., 2019) and
71 palaeo-hydroclimatic variability (Passey et al., 2014; Passey and Ji, 2019; Bergel et al., 2020;
72 Huth et al., 2022; Sha et al., 2020, 2023, 2024a). However, the isobaric interference of
73 $^{13}\text{C}^{16}\text{O}^{16}\text{O}$ with $^{12}\text{C}^{17}\text{O}^{16}\text{O}$ necessitates an exceptionally high mass resolving power ($>56,000$).
74 Even at such resolution, achieving high-precision measurements of ^{17}O in CO_2 is challenging
75 due to its low natural abundance and signal suppression by the intense tail of the $^{13}\text{C}^{16}\text{O}^{16}\text{O}$
76 peak. This technical challenge has spurred the development of several alternative techniques
77 that either produce O_2 from CO_2 or transfer the oxygen isotopic signature of CO_2 to O_2 . These
78 include fluorination by BrF_5 , methanation-fluorination, CO_2 - H_2O exchange-fluorination, CO_2 -
79 CeO_2 exchange-laser fluorination and Pt catalysed CO_2 - O_2 exchange (Bhattacharya and
80 Thiemens, 1989; Hofmann and Pack, 2010; Barkan and Luz, 2012; Mahata et al., 2013; Passey
81 et al., 2014). The CO_2 - O_2 exchange method, which has been first introduced by Mahata et al.,
82 2013 with subsequent modifications by Barkan et al., 2015, remains the most adopted
83 technique due to its clean and simple approach with the best available measurement precision.
84 Mahata et al., 2013 determined the $\delta^{17}\text{O}$ of CO_2 by measuring the $\delta^{18}\text{O}$ of CO_2 before exchange
85 and $\delta^{18}\text{O}$ and $\delta^{17}\text{O}$ of O_2 before and after exchange. For that, they used the slope of the line
86 connecting the initial and final isotope ratios and applied a minor correction based on the
87 deviation of the final ratios from this line. Presently, $\delta^{17}\text{O}$ of CO_2 is mostly measured using the
88 calibrated isotope exchange fractionation factors ($^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$). The ratio of natural
89 logarithms of the two oxygen isotope fractionation factors is termed as the triple oxygen isotope
90 exponent ($\theta_{\text{CO}_2/\text{O}_2}$). The deviation of $\delta^{17}\text{O}$ from that expected for a given $\delta^{18}\text{O}$, based on a
91 reference slope, is referred to as the triple oxygen isotope anomaly ($\Delta^{17}\text{O}$). The reference slope
92 is denoted as λ when multiple fractionation processes are involved and as θ when the
93 fractionation process is defined (Passey et al., 2014; Miller and Pack, 2021b; Aron et al., 2021).

$$94 \theta_{\text{CO}_2/\text{O}_2} = \frac{\ln ^{17}\alpha_{\text{CO}_2/\text{O}_2}}{\ln ^{18}\alpha_{\text{CO}_2/\text{O}_2}}$$

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$$96 \quad \Delta^{17}\text{O} = \delta^{17}\text{O} - \lambda_{\text{reference}} \times \delta^{18}\text{O}; \delta' = \ln \left(\frac{R_{\text{sample}}}{R_{\text{standard}}} \right)$$

97 The temperature at which the exchange is carried out by most researchers is ~ 750 °C. The
98 $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}$, at this temperature, derived from internal partition function ratios
99 using CO₂ datasets from Tashkun and Harvey, 2025 and our oxygen datasets, are 1.004390 and
100 1.002323, respectively (discussed in section 2.4). Replacing CO₂ datasets with that of Huang
101 et al., 2017, the corresponding values become 1.004490 and 1.002377. The slight differences
102 in the calculated values arise primarily from the uncertainty in the zero-point energy (ZPE)
103 differences between CO₂ isotopologues. But the experimentally determined $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$,
104 $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$, at 750 °C, are far from theory and vary significantly across laboratories,
105 ranging from 1.00051 to 1.00227, 1.00048 to 1.00135 and 0.5 to 0.9, respectively (Barkan et
106 al., 2015; Fosu et al., 2020; Adnew et al., 2022). As the use of the proxy advances with
107 expansion into new fields, establishing high-precision, consistent inter-laboratory values has
108 become increasingly critical. Deviations from the equilibrium conditions were originally
109 thought to derive from heterogeneous processes, but Adnew et al., 2022 ruled out this
110 possibility and suggested that thermal diffusion effects cause these deviations. Following up
111 on this proposition, Wei et al., 2024 conducted thermal-diffusion experiments on pure O₂, pure
112 CO₂ and their mixtures without Pt to isolate thermal gradient effects from exchange reactions.
113 They discussed how $\Delta\delta^{18}\text{O}_{\text{hot-cold}}$ and fraction of gas in the hot zone influence the fractionation
114 factors of Pt-catalysed exchange reactions based on their thermal diffusion experiments.
115 However, in their experiments, as only half of the designated hot zone was heated in the
116 furnace, a thermal gradient must have existed within the presumed hot zone itself. The
117 assumption of a simple two-zone (hot and cold) model along the entire extraction line is
118 oversimplified and may have contributed more to the observed discrepancies in $\theta_{\text{O}_2}^{\text{TG}}$
119 (fractionation exponent of thermal diffusion for O₂) than a different temperature range in the
120 previous study.

121 In the present study, we introduce modifications to the existing methods, incorporating key
122 innovations to estimate $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values across reactors with varying cold zone, ranging from
123 78.9 to 0%. A key challenge addressed in this study is exchanging CO₂-O₂ in complete absence
124 of thermal gradient effects, which is essential to confirm that any deviation from the theoretical
125 $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values arise only from such effects. As thermal gradients are unavoidable in reactors
126 used by most laboratories for $\Delta^{17}\text{O}$ measurements in CO₂, it is critical to develop a framework

127 for standardizing triple oxygen isotope measurements under varying proportions of hot and
128 cold zones to avoid interlaboratory inconsistencies. The specific objectives of this study are to:
129 (i) accurately determine $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values while progressively minimising the
130 thermal gradient effects to nil, (ii) experimentally verify the underlying reasons for large
131 deviations from theoretical predictions and (iii) recommend appropriate $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values to
132 establish a standardized method for triple oxygen isotope measurements in CO_2 .

133

134 **2. Materials and methods**

135 **2.1 Experimental setup**

136 Figure 1 illustrates the in-house high vacuum setup designed for triple oxygen isotope
137 measurements. High-purity CO_2 and O_2 gases (99.999%), sourced from ATCO Atmospheric
138 and Speciality Gases Private Limited, Ahmedabad, are introduced into the system via stainless
139 steel tubing. The gas pressure during transfer is regulated using Setra™ pressure transducers
140 connected to the vacuum system. The system's main component, the exchange-reactor, consists
141 of two sections separated by a high vacuum valve, necessary for sequential transfer of the two
142 gases before exchange. The CO_2 extracted from various sources such as carbonates, air and
143 water- CO_2 equilibration experiments is cryo-trapped in the lower section of the exchange
144 reactor. An approximately equal amount of O_2 is expanded from an isotopically known tank
145 into the upper section keeping the middle valve closed. The two gases are mixed and heated at
146 the desired temperature keeping only the lower section inside the heater. Heating is performed
147 using Watlow™ ceramic fibre heaters, controlled through an Autonics™ temperature controller
148 and monitored in real time using LabView software.

149 **2.2 CO_2 - O_2 exchange experiments under different proportions of hot and cold zones**

150 In order to study the influence of the thermal gradient on the exchange equilibrium, a series of
151 experiments with varying proportions of hot and cold zones have been performed. For all
152 experiments, a 1:1 mixture of CO_2 and O_2 gases with known values of isotopic composition
153 was prepared ($\delta^{18}\text{O}_{\text{O}_2} = 26.399 \text{ ‰}$, $\delta^{17}\text{O}_{\text{O}_2} = 13.332 \text{ ‰}$, $\delta^{18}\text{O}_{\text{CO}_2} = 1.185 \text{ ‰}$ and $\delta^{17}\text{O}_{\text{CO}_2} =$
154 0.593 ‰). The isotopic compositions of these CO_2 and O_2 were measured at Academia Sinica,
155 Taiwan. CO_2 was transferred into an evacuated glass flask (to a pressure of ~ 250 torr), after
156 which O_2 was carefully introduced from the oxygen cylinder through the vacuum line until the
157 total pressure doubled (~ 500 torr), resulting in equal proportions of CO_2 and O_2 in the flask.
158 The O_2 pressure outside the glass flask was maintained at a significantly higher level to prevent

159 back-diffusion of CO₂ from the flask. The oxygen isotope fractionation factors of exchange,
160 $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$, and triple oxygen exponent $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ were determined at four
161 temperatures across the range of 600 to 1000 °C. Exchange reactors with four different
162 configurations were designed to modify the volume percent remaining outside the heater
163 (referred to as the cold zone hereafter). In the first set of experiments, the mixture was heated
164 in the presence of 150 mg Pt-sponge in the exchange reactor with 78.9% cold zone for 1.5
165 hours. It has been previously reported through controlled experiments that $\delta^{18}\text{O}$ and $\delta^{17}\text{O}$ may
166 take longer than an hour to stabilise to the steady-state values within 0.01-0.02% (Mahata et
167 al., 2013; Adnew et al., 2022; Barkan et al., 2015). For the second set, the CO₂-O₂ mixture was
168 exchanged over Pt only in the lower part of the reactor keeping the middle valve closed (see
169 Fig. 1), thereby restricting the cold zone to 62.2%. Further, a special ‘reduced cold zone’ reactor
170 was devised to constrain the cold zone to 45%. There was no middle valve separating this
171 reactor into two sections, limiting its use to heating a CO₂-O₂ mixture and making it unsuitable
172 for routine sample measurements. We attempted to further reduce the cold zone of this modified
173 reactor to 25% and exchange reactions were conducted at ~700 °C. However, the borosilicate-
174 quartz seal, due to its thermal sensitivity, failed during high-temperature exchange. Also, a few
175 initial high-temperature exchange experiments produced inconsistent results, likely because
176 the Pt at the base was not thermally stable due to the heater-reactor geometry.

177 During collection after exchange, CO₂ was first trapped in the U-trap (Fig. 1) in liquid nitrogen
178 and O₂ was collected on silica pellets dipped in liquid nitrogen for about 5 minutes. Afterwards,
179 CO₂ was collected from the U-trap. All exchange experiments at each reactor temperature were
180 conducted in at least triplicate. Each sample was measured over five cycles, with signal
181 intensities ranging from 3000 to 4000 mV for O₂ (mass 32) and 6000-8000 mV for CO₂ (mass
182 44). Measurements were conducted with an integration time of 10 seconds and an idle time of
183 24 seconds between cycles. The reported errors represent the 1- σ standard deviations from
184 replicate measurements.

185 To completely eliminate the cold zone, an innovative approach was adopted. The CO₂-O₂
186 mixture together with conditioned Pt was sealed in quartz tubes which were heated inside a
187 muffle furnace at 650 °C, 700 °C, 750 °C and 800 °C. The furnace ensured homogeneous
188 heating of the tubes forming ‘no cold zone’ reactors. After heating, the sealed tubes were
189 immediately quenched in water to halt further isotopic exchange, as it is well established that
190 the two gases do not exchange under normal conditions (Katakis and Taube, 1962). The
191 exchanged gases were recollected using a break-seal arrangement in the same vacuum-setup

192 and analysed, following the procedure described above. To validate the attainment of
 193 equilibrium in these experiments, we performed another experiment in which the mixture was
 194 heated to 650 °C for two hours and then the temperature was increased to 800 °C for another
 195 two hours. After exchange with O₂ at 650 °C, the CO₂ would have attained a higher δ¹⁸O of
 196 17.51 ‰ (Table S2) from its original value of 1.18 ‰ as for the normal heating at 650 °C. Now
 197 when the temperature is increased to 800 °C where the fractionation factors are lower, the δ¹⁸O
 198 of CO₂ will have to decrease from its original value to reach the steady-state value rather than
 199 increasing from the usual 1.18 ‰. Exchange experiments were limited to temperatures ≤ 800
 200 °C owing to furnace instability at higher temperatures.

201 **2.3 Determining the oxygen fractionation factors for CO₂-O₂ exchange**

202 The isotopic compositions of the exchanged gases were analysed using a stable isotope ratio
 203 mass spectrometer (MAT 253) in dual-inlet mode at Physical Research Laboratory,
 204 Ahmedabad. The δ¹⁸O and δ¹⁷O of O₂ were measured directly. The final δ¹⁸O and δ¹³C of the
 205 CO₂ were measured with an internal precision of 0.05‰. The δ¹³C of CO₂ served as a control
 206 to identify any fractionation that might have occurred during sample collection, as it is expected
 207 to remain constant throughout the exchange. The heaters employed, however, had relatively
 208 higher temperature fluctuation of ±5 °C, resulting in added errors across repeated batches. The
 209 β values, the CO₂/O₂ molar ratio (Mahata et al., 2013; Barkan et al., 2015), calculated using
 210 oxygen mass balance during exchange are close to 1 as expected for a mixture prepared with
 211 approximately equal amounts of CO₂ and O₂.

$$212 \quad \beta = \frac{\delta^{18}\text{O}_{\text{Initial}}(\text{O}_2) - \delta^{18}\text{O}_{\text{Final}}(\text{O}_2)}{\delta^{18}\text{O}_{\text{Final}}(\text{CO}_2) - \delta^{18}\text{O}_{\text{Initial}}(\text{CO}_2)}$$

213 The ¹⁸α<sub>CO₂/O₂ for exchange could be directly calculated while ¹⁷α<sub>CO₂/O₂ was derived from the
 214 known initial δ¹⁷O of CO₂ using procedure described in Barkan et al. (Barkan et al., 2015).</sub></sub>

$$215 \quad {}^{17}\alpha = \frac{1}{\beta} \left\{ \left[\frac{\delta^{17}\text{O}_{\text{Initial}}(\text{CO}_2) + 1}{[(\delta^{17}\text{O}_{\text{Final}}(\text{O}_2) + 1)]} \right] - 1 \right\}$$

216

217 **2.4 Theoretical estimation of ¹⁸α_{CO₂/O₂ and ¹⁷α_{CO₂/O₂}}**

218 Thermodynamic equilibria for the oxygen isotope exchange between gaseous O₂ and CO₂ is
 219 expressed as



221 where i indicates a heavy isotope with mass number 17 or 18, have been calculated according
 222 to statistical theory, pioneered by Bigeleisen and Mayer, 1947 and Urey, 1947 (BMU,
 223 hereafter). Here, we refine our previous calculations (Adnew et al., 2022; Prokhorov et al.,
 224 2019). The fractionation factor for the above isotope exchange equilibrium reaction is given by
 225 (Prokhorov et al., 2019; Adnew et al., 2022). The fractionation factor for the above isotope
 226 exchange equilibrium reaction is given by

227
$${}^i\alpha_{\text{CO}_2/\text{O}_2} = \frac{Q(\text{O}_2)}{Q({}^i\text{OO})} \frac{Q(\text{C}^i\text{OO})}{Q(\text{CO}_2)} =$$

228
$$\left(\frac{M(\text{O}_2)M(\text{C}^i\text{OO})}{M({}^i\text{OO})M(\text{CO}_2)}\right)^{3/2} \frac{Q_{int}(\text{O}_2)}{Q_{int}({}^i\text{OO})} \frac{Q_{int}(\text{C}^i\text{OO})}{Q_{int}(\text{CO}_2)} e^{-\left(\frac{\epsilon_0(\text{O}_2) - \epsilon_0({}^i\text{OO}) + \epsilon_0(\text{C}^i\text{OO}) - \epsilon_0(\text{CO}_2)}{kT}\right)} \quad (\text{A})$$

229 In this equation, M denote molecular masses and ϵ_0 the ZPEs, whereas Q and Q_{int} stand for
 230 total and internal partition functions, respectively.

231
$$Q_{int} = \sum_i d_i e^{-\frac{\epsilon_i}{kT}} \quad (\text{B})$$

232 The internal partition function Q_{int} is the sum over all internal states viz, ro-vibrational and
 233 electronic states i with state degeneracy d_i and energy ϵ_i above ϵ_0 . We note in passing that the
 234 relative uncertainty of ${}^i\alpha_{\text{CO}_2/\text{O}_2}$ is factually the quadratic sum of four independent contributions:
 235 the relative uncertainties of the two internal partition function ratios $Q_{int}(\text{C}^i\text{OO})/Q_{int}(\text{CO}_2)$
 236 and $Q_{int}({}^i\text{OO})/Q_{int}(\text{O}_2)$ and the relative uncertainties of the ZPE difference terms
 237 $\exp((\epsilon_0(\text{O}_2) - \epsilon_0({}^i\text{OO}))/kT)$ and $\exp((-\epsilon_0(\text{CO}_2) + \epsilon_0(\text{C}^i\text{OO}))/kT)$. The contributions from
 238 atomic and molecular masses in eq. (A) can be safely neglected due to their high accuracy.

239 Electronic excited states will play only a small role in the evaluation of the partition functions
 240 and they are just considered in the case of O_2 . For di-oxygen, these excited states have adiabatic
 241 energies of ~ 1 eV (7918 cm^{-1}) or more above the ground electronic state and, compared to ro-
 242 vibrational states in the ground state, those in the first excited electronic state have therefore
 243 an additional relative weight of 1.3×10^{-4} at $1000 \text{ }^\circ\text{C}$. The effect on the ratio of partition
 244 functions of two isotopologues of the same molecule is smaller, because these have similar
 245 state energies on the same potential energy surface and error cancellation occurs when dividing
 246 the two similar partition functions. The same arguments apply even more to CO_2 with the
 247 lowest excited electronic state energies being ~ 4 to 5 eV (Spielfiedel et al., 1992; Ma et al.,
 248 2014). At these energies, the Boltzmann factor in eq. (B) becomes as low as 10^{-16} at $1000 \text{ }^\circ\text{C}$,

249 and these states can thus be completely neglected in the evaluation. If the energies of all states
250 up to a sufficiently high threshold are known to spectroscopic accuracy, one can directly
251 evaluate the above sums (eq. B). We have done so using the CO₂ data provided by Huang et
252 al., 2017, which are based on theoretical molecular calculations, improved by comparison with
253 spectroscopic observations. Note, however, that due to inadvertently using $d_J = J$ instead of d_J
254 $= (2J+1)$ as rotational degeneracy factors in the calculation of partition sums of CO₂
255 (unpublished work), Adnew et al., 2022 have obtained slightly different results than we do here.
256 Internal partition function ratios, such as $Q_{int}(C^iOO)/Q_{int}(CO_2)$ from our calculation agree
257 with a former direct sum evaluation (unpublished work) for relevant temperatures above
258 500 °C to $\lesssim 10^{-4}$ relative, which is close to rounding errors of $\sim 5 \times 10^{-5}$. Moreover, there is even
259 better agreement with the newly recommended values of $Q_{int}(C^iOO)/Q_{int}(CO_2)$ from
260 Tashkun and Harvey, 2025, where differences are smaller than 2×10^{-5} over the temperature
261 range between 100 and 1300 K. However, the main uncertainty of α due to CO₂ stems from the
262 ZPE difference terms $\epsilon_0(C^iOO) - \epsilon_0(CO_2)$. The two data sets^{22,23}, yield ZPE differences that
263 differ by 0.06 cm⁻¹ and 0.03 cm⁻¹ for ¹⁸O and ¹⁷O containing isotopologues, respectively. This
264 leads to deviations between the two calculations of about 1.1×10^{-4} (for ¹⁸O) and 0.6×10^{-4} (for
265 ¹⁷O) at low (800 K), and 6.7×10^{-5} and 3.6×10^{-5} at high (1300 K) temperatures. Without any
266 further a-priori knowledge, we use the values at 800 K derived from the two completely
267 independent theoretical approaches to estimate relative uncertainties $u_r(Q_{int}(C^iOO)/$
268 $Q_{int}(CO_2)) = 0.11 \%$ and 0.06% , for ¹⁸O and ¹⁷O respectively.

269 For the partition functions of O₂, we have previously (Huang et al., 2017; Janssen and Tuzson,
270 2010) used the analytical BMU approach, even though it has not yet been demonstrated
271 whether the approach can actually be applied correctly to O₂. While the BMU method has
272 originally been developed for closed shell molecules, its direct application to molecular oxygen
273 neglects that the molecule has an electronic fine structure due to its non-zero total electron spin,
274 which includes spin-spin and spin-rotation terms (Yu et al., 2012). Even though these fine
275 structure couplings are quite sizable on the order of a few cm⁻¹, they have only a very weak
276 isotope dependence and could possibly cancel in the calculation of partition function ratios. In
277 the case of oxygen, however, quantum mechanical symmetry rules restrict allowed states of the
278 symmetric ¹⁶O₂ isotopologue to odd quantum numbers N for the molecular rotation, while odd
279 and even rotational states of the asymmetric isotopologues ¹⁷OO and ¹⁸OO are allowed. This
280 implies that the rotational ground state of the homonuclear ¹⁶O₂ molecule has a higher quantum
281 number ($N = 1$) and rotational energy than the heteronuclear molecules in their $N = 0$ states. As

282 a consequence, one obtains an anomalous ZPE scaling with a difference ratio of
283 $(\epsilon_0(^{17}\text{O}_2) - \epsilon_0(\text{O}_2)) / (\epsilon_0(^{18}\text{O}_2) - \epsilon_0(\text{O}_2)) = (-12.379 \text{ cm}^{-1}) / (-22.869 \text{ cm}^{-1}) = 0.541$, as
284 compared to $(-11.669 \text{ cm}^{-1}) / (-22.145 \text{ cm}^{-1}) = 0.527$ when fine structure effects and symmetry
285 restrictions are ignored. Previous calculations that neglected electron spin and symmetry
286 restrictions on the ZPEs therefore must be subject to systematic error, which should become
287 most apparent at low temperatures. In order to avoid such bias, we have included electronic
288 fine-structure effects in our calculation. The details of the calculation will be presented
289 elsewhere (unpublished work). Here, we just summarise the most salient conclusions, and
290 report on the uncertainty estimate in the relevant temperature range between 800 and 1300 K
291 (or about 500 to 1000 °C). The calculations are based on a global evaluation of spectroscopic
292 constants of all O₂ isotopologues made by Yu et al., 2014, 2012, which includes all three lowest
293 electronic states with vibrational states up to $v = 35$ for the ground electronic state and $v = 17$
294 for the two lowest excited states. Energy levels for all stable isotopologues are provided in the
295 supplementary material of the study (Yu et al., 2014), but are restricted to rotational quantum
296 numbers N smaller than 66, roughly corresponding to energies below 6080 cm^{-1} . This is too
297 small to accurately calculate partition functions for temperatures up to 1300 K. We have
298 therefore added energies for missing ro-vibrational states in the electronic ground state by
299 extrapolating term energies towards higher rotational excitation using the spectroscopic
300 constants provided by Yu et al., 2014. The validity of this approach has been checked by
301 comparing our calculated internal partition functions for all six isotopologues of O₂ to the
302 results presented in Table 1 of Yu et al., 2012 which cover the temperature range between 9.375
303 and 300 K, and to the very recent internal partition function of ¹⁶O₂ provided by Furtenbacher
304 et al., 2025. Based on an even larger set of spectroscopic data and more excited electronic
305 states, the latter covers the temperature range between 1 and 5000 K and also specifies
306 uncertainties as a function of temperature. Our calculations agree with the results of Yu et al.,
307 2014 and our agreement with (Furtenbacher et al., 2025) under 1300 K is always better than
308 5×10^{-7} . The relative standard uncertainty of $Q_{int}(^{16}\text{O}_2)$ indicated by (Furtenbacher et al., 2025)
309 for this temperature range is 2×10^{-5} or smaller. Even their highest value is smaller than the
310 uncertainties that we have determined in the evaluation of CO₂. The uncertainty of the ZPE
311 differences can be conservatively estimated using the evaluation of ZPEs and their uncertainties
312 from (Irikura, 2007). Based on a smaller set of spectroscopic data and using a less complete
313 and more ancient set of molecular constants than Yu et al., 2014, Irikura, 2007 obtains a
314 standard uncertainty of $4.5 \times 10^{-3} \text{ cm}^{-1}$ for the ZPE of ¹⁶O₂ when the Dunham terms up to the

315 sixth order are available. $\epsilon_0(^{18}\text{O}) - \epsilon_0(^{16}\text{O})$ is thus accurate to 0.01 cm^{-1} or better, thus leading
316 to an uncertainty of less than 2×10^{-5} at temperatures as low as $500 \text{ }^\circ\text{C}$.

317 To summarize, in the $500 \text{ }^\circ\text{C}$ to $1000 \text{ }^\circ\text{C}$ range, the uncertainties of $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}$ are
318 dominated by the uncertainty in the ZPE differences between CO_2 isotopologues. They lead to
319 amount to standard uncertainties of $0.11 \text{ }^\circ\text{‰}$ and $0.06 \text{ }^\circ\text{‰}$ for $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}$,
320 respectively.

321 3. Results

322 3.1 Variations in $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ with varying cold zone volumes

323 Measured $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values with varying cold zone volumes are presented in
324 Figure 2 and Supplementary Information Table S2. The values are compared with the
325 theoretically predicted equilibrium values. When no cold zone was present, i.e., the entire CO_2 -
326 O_2 mixture was uniformly heated inside the furnace, the measured $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$
327 matched the theoretical predictions within experimental uncertainty. Although the experimental
328 values at all temperatures are marginally higher than theoretical estimates, the deviations are
329 insignificant. To confirm that this small offset was not due to incomplete equilibration, an
330 additional experiment was conducted in which the mixture was first equilibrated at $650 \text{ }^\circ\text{C}$ and
331 subsequently heated to $800 \text{ }^\circ\text{C}$. The rationale was that at $650 \text{ }^\circ\text{C}$, $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ would
332 attain relatively higher equilibrium values, and subsequent heating at $800 \text{ }^\circ\text{C}$ would lower them
333 again. The consistency between this two-step experiment and direct equilibration at $800 \text{ }^\circ\text{C}$
334 confirmed attainment of complete isotopic equilibrium (Fig. 2). The potential causes of the
335 minor systematic deviation have been detailed in the section 4.3. As the hot zone volume
336 decreased, the deviation from the theoretical $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}$ values became more
337 pronounced. For instance, at a cold zone proportion of approximately 45%, the deviations were
338 ~ 0.004 for $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ and ~ 0.002 for $^{17}\alpha_{\text{CO}_2/\text{O}_2}$ at $750 \text{ }^\circ\text{C}$. These discrepancies continued to
339 grow with further increase in cold zone volume, indicating a progressive departure from
340 equilibrium conditions. The $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ obtained from the limited exchange done
341 at $\sim 700 \text{ }^\circ\text{C}$ keeping 25% cold zone still deviated from the corresponding theoretical predictions
342 by 0.002 and 0.001 (Supplementary Information Table S4).

343 An important observation related to exchanges involving a cold zone is the behaviour of the
344 fractionation factors $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$. These values are greater than 1 at lower

345 temperatures, but drop below 1 once the temperature exceeds a certain threshold (Fig. 2 and
346 3). Notably, the specific temperature at which this transition occurs depends on the fraction of
347 the cold zone. Similar trends were reported by Adnew et al., 2022 and predicted by the model
348 in Wei et al., 2024 . In setups without a cold zone, the α values consistently remain above 1 and
349 close to the theoretical equilibrium values, gradually approaching 1 as the temperature
350 increases (Fig. 2 and 3). The transition behaviour has been attributed to thermal diffusion
351 effects between the hot and cold regions in setups with a cold zone.

352 Another intriguing observation is the crossover of α values: at lower exchange temperatures,
353 $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ is lower than $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$, but at higher temperatures, this trend reverses (Fig. 3). The
354 temperature at which this crossover occurs also depends on the fraction of the hot zone. This
355 crossover is not observed in the absence of a cold zone, consistent with theoretical expectations.
356 Although Adnew et al., 2022 reported that the crossover coincides with the exchange-inflection
357 temperature (temperature at which α transits from >1 to <1), our results show that this is not
358 always the case (Fig. 3). In our experiments, the temperature at which the crossover occurs
359 does not necessarily align with the $\alpha = 1$ transition point (Fig. 3).

360 **3.2 Comparing $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ among reactors with varying cold zones**

361 The average $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ value of $0.529 (\pm 0.015)$, calculated from the sealed quartz tubes in the
362 temperature range 650 to 800°C, agree well with the theoretical value of 0.529 (Fig. 4,
363 Supplementary Information Table S1 and S2). Conceptually, θ is an intrinsic property of a
364 defined process (Bao et al., 2016). The reaction occurring in conventional exchange reactor is
365 a combination of isotope-exchange and thermal diffusion leading to a large variability in
366 $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values (Fig. 4). These values are much beyond the canonical range in exchange
367 reactions conducted at or above the crossover temperature associated with the respective cold
368 zone. Wei et al., 2024 predicted a relationship between $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ which is included
369 in Figure 4 using our revised equilibrium $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ value and $\theta_{\text{O}_2}^{\text{TG}}$ from their diffusion
370 experiments. Although negative $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ have been predicted under conditions of extreme
371 thermal diffusion by their model, increasing cold zones do not produce such behaviour. An
372 exception is the $\sim 880^\circ\text{C}$ experiment with 45% cold zone. The large deviation observed for this
373 data arises from the diverging behaviours of $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ (by definition) under these conditions:
374 $\ln^{17}\alpha_{\text{CO}_2/\text{O}_2}$ approaches zero while $\ln^{18}\alpha_{\text{CO}_2/\text{O}_2}$ becomes negative (Fig. 4) leading to an
375 amplified and unstable $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ value.

376 4. Discussion

377 4.1 Experimental versus theoretical $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$

378 Theoretically $\alpha_{\text{CO}_2/\text{O}_2}$ values for exchange are expected to be larger than 1 and approach unity
379 at very high temperature (Richet et al., 1977; Adnew et al., 2022). Moreover, $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ should
380 always be greater than $^{17}\alpha_{\text{CO}_2/\text{O}_2}$. However, in the experimental exchange done in reactors with
381 varying volumes outside the heater, the $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values are not always greater than 1 due to
382 thermal diffusion between hot and cold zones. The $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values are higher than 1 at lower
383 temperatures and decrease to less than 1 above a certain temperature in all three reactors with
384 78.9%, 62.2% and 45% cold zones. The temperature of this exchange-inflection point increased
385 with decreasing cold zones (Fig. 3). Fractionation due to thermal-gradient can explain the
386 selective enrichment of O₂ over CO₂ after the exchange-inflection point. Wei et al., 2024
387 showed that oxygen isotope fractionation of O₂ is greater than that of CO₂ when subjected to
388 identical thermal gradient conditions and was attributed to isotopic fractionation due to thermal
389 diffusion. The latter topic has been discussed in details by Jones and Furry, 1946. The O₂ in the
390 cold zone of the reactor is more enriched in heavy isotopes compared to CO₂, resulting in values
391 of $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ lower than 1.

392 A cross over temperature in $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values was also observed: below this temperature,
393 $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ is greater than $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ whereas above it, $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ exceeds $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$. This
394 inversion, where the abundance of ¹⁷O surpasses that of ¹⁸O in both CO₂ and O₂, suggests the
395 involvement of a kinetic effect during the exchange process caused by thermal gradient
396 between the cold and hot zones. Such thermal gradient caused a systematic deviation from the
397 isotopic equilibrium, even though thermodynamic equilibrium was attained after 0.5 to 1 hour
398 of heating. All other previous studies (e.g. Mahata et al., 2013; Barkan et al., 2015; Adnew et
399 al., 2022; Wei et al., 2024) that employed setups with distinct hot and cold zones also likely
400 observed thermodynamic equilibrium states that were significantly offset isotopically due to
401 thermal diffusion. The thermal gradient induced kinetic effect was supported by clumped
402 isotope measurements in CO₂ and O₂ following heating of a CO₂ - O₂ mixture in a typical setup
403 with both hot and cold zones (unpublished data; personal communication, Mao-Chang Liang),
404 because $\Delta_{47}(\text{CO}_2)$ and $\Delta_{36}(\text{O}_2)$ values were far from equilibrium. In our system, the crossover
405 temperature does not coincide with the exchange-inflection temperature. The overlap of these
406 two points observed by Adnew et al., 2022 is likely coincidental and may result from the
407 specific geometry of their experimental setup. Notably, the crossover is not observed in

408 exchanges conducted within sealed quartz tubes heated uniformly. These experiments attained
409 isotopic equilibrium, showing close agreement with theoretical predictions. The minor
410 discrepancies observed between the theoretical values and those from no-cold zone
411 experiments fall within the expected uncertainty of the experimental values. The experiments
412 in which the gas mixtures were first heated to 650°C for two hours and subsequently at 800 °C
413 for another two hours yielded $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values indistinguishable from those obtained in
414 experiments heated directly at 800°C, further confirming the attainment of complete
415 equilibrium under uniform heating conditions.

416 At 750 °C, the offsets in $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ determined with 0%, 45%, 62.2% and 78.9% cold zones are
417 0.0003, 0.004, 0.004 and 0.005, respectively, from the theoretical value of 1.00449. This
418 implies that reducing cold zone volumes in the exchange reactor alone may not be sufficient to
419 achieve precise triple oxygen measurements in carbonates. Moreover, the reactor configuration
420 makes it challenging to further minimise the cold zone in routine measurements. Interestingly,
421 the results show a strong dependence of $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ on hot zone volume, offering a potential
422 pathway for standardizing the exchange reactions.

423 **4.2 Relation between $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and reactor configuration**

424 An exponential relationship between the fractionation factors and the hot zone volumes has
425 been established at 750°C (Fig. 5). Although the exact hot and cold zone volumes are subject
426 to some uncertainty due to the irregular geometry of the reactor tubes, the observed trend
427 remains significant. Further, to test the validity of fractionation factors derived from the
428 exponential relation in figure 5, an isotopically different CO₂ was exchanged with the same O₂.
429 The $\delta^{18}\text{O}$ value of this new CO₂ was 37.23 ‰ in contrast to the original CO₂ having $\delta^{18}\text{O}$ value
430 of 1.18 ‰. It is to be noted that the $\delta^{18}\text{O}$ value of the new CO₂ was higher than that of O₂ (26.40
431 ‰), unlike the previous case. The exchange was carried out in two different reactors with
432 41.9% and 62.2% cold zones. The measured $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ were in good agreement with those
433 expected from the exponential fit (Supplementary Information Table S3). This exponential
434 relationship was also tested at 700°C including fractionation factors from the additional
435 exchange reaction conducted with 25% cold zone proportion (Supplementary Information
436 Figure S1 and Table S4). The observed trend was significant and provided further validation of
437 the proposed framework.

438 In practical setups, complete elimination of cold zone is not feasible. Therefore, it is essential
439 to calibrate triple oxygen isotope exchange setups based on a fixed configuration of hot and
440 cold zones, and accurately determine the associated fractionation factors. For setups similar to
441 ours, where the hot and cold zone volumes are constrained, the recommended fractionation
442 factors reported here may be applicable. However, before adopting them as standardized
443 factors, this should be verified experimentally by independent groups. We also plotted selected
444 $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values from previous studies that employed similar experimental setups (Fig. 5). Due
445 to the lack of detailed information regarding the hot and cold zone proportions in those studies,
446 we assumed a 50:50 ratio between the heated and cold sections of the reactors in all the cases.
447 Most of these values showed significant deviations from our fitted curve, with the exception of
448 the data from Adnew et al., 2022 in which the authors explicitly stated that half of the reactor
449 was within the heating zone and results are in good agreement with our fitted line. The
450 discrepancies observed for other studies highlight the importance of accurately characterizing
451 the thermal geometry of the reactor when interpreting fractionation factors.

452 **4.3 Limitations**

453 We observed a small but systematic deviation from the theoretical fractionation values, as
454 discussed earlier. This discrepancy likely stems from a combination of factors, including (i)
455 uncertainties in the theoretical zero-point energy calculations (though these are expected to be
456 small, on the order of ~ 0.0005), (ii) experimental errors such as potential biases in the
457 $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ and $^{17}\alpha_{\text{CO}_2/\text{O}_2}$ values of the CO_2 and O_2 used for exchange, and (iii) slight
458 inaccuracies in assigning the actual experimental temperature. Liang et al., 2023 reported a bias
459 of ~ 0.001 in $\theta_{\text{CO}_2/\text{O}_2}$ between Academia Sinica and the Hebrew University of Jerusalem, and a
460 difference of $\sim 0.04\%$ in $\Delta^{17}\text{O}$. As our CO_2 and O_2 were calibrated in Academia Sinica, small
461 offsets in the $\delta^{17}\text{O}$ and $\delta^{18}\text{O}$ values of the exchange gases could contribute to the deviations
462 from theoretical expectations. Temperature assignment introduces another potential source of
463 bias. The exchange temperature was taken as that measured when the sealed quartz tube was
464 inside the furnace. However, once the furnace is opened and the tube is rapidly quenched in
465 water, brief partial exchange of the CO_2 - O_2 mixture at lower temperatures cannot be completely
466 ruled out. This could introduce a consistent temperature-related offset across all experiments.
467 An exchange occurring at temperatures 30–50 °C lower than assumed would be sufficient to
468 explain the observed discrepancies between theoretical and experimental values.
469 Unfortunately, estimating the actual effective re-equilibration temperature for these setups

470 remains challenging. Further experiments conducted independently by other laboratories will
471 be essential to refine the exchange fractionation factors and improve the robustness of triple
472 oxygen isotope measurements in CO₂.

473 **4.3 Broader implications for triple oxygen isotopes measurements in CO₂**

474 ¹⁷O has emerged as a promising proxy that complements ¹⁸O in tracing the global hydrological
475 cycle and reconstructing palaeo-environments, particularly in archives where d-excess is not
476 commonly available. Recent studies have used Δ¹⁷O in carbonates to quantify past relative
477 humidity conditions, for example, parent water derived from speleothems (Sha et al., 2020,
478 2023) and palaeo-seawater reconstructions from foraminifers (Sha et al., 2024b). These
479 investigations report sensitivities of approximately 1 permeg Δ¹⁷O per % relative humidity
480 change. Such levels of sensitivity underscore the need for highly precise and inter-laboratory
481 comparable Δ¹⁷O measurements. However, reported Δ¹⁷O of international carbonate
482 standards show measurable discrepancies across laboratories (Supplementary Information
483 Table S5 and S6). For instance, Δ¹⁷O values of CO₂ derived from NBS-18 span a range of
484 ~106 permeg across studies employing Pt-catalysed exchange method (Liang et al., 2017a; Sha
485 et al., 2020; Fosu et al., 2020; Barkan et al., 2019) and are much lower than those obtained
486 using the reduction-fluorination method (Passey et al., 2014; Passey and Ji, 2019; Wostbrock
487 et al., 2020; Huth et al., 2022). Differences in acid digestion temperature between the two
488 methods were initially proposed as a potential source of this discrepancy. However, Wostbrock
489 et al., 2020 performed acid digestion at 25°C followed by reduction-fluorination and reported
490 values consistent with those obtained at 90°C by Passey et al., 2014. We suspect that the
491 significantly lower Δ¹⁷O values measured for NBS-18, NBS-19 and IAEA-603 using Pt-
492 catalysed exchange than reduction-fluorination method may partly reflect kinetic effects
493 induced by thermal gradient. Consequently, consensus Δ¹⁷O values based on the Pt-catalysed
494 exchange remain ambiguous, as the associated exchange fractionation factors are neither
495 consistent amongst laboratories nor aligned with theoretical predictions. Our novel approach
496 that allows exchange under true equilibrium (0% cold-zone) and has yielded fractionation
497 factors and exponents consistent with theory, can be utilized to verify Δ¹⁷O values of the
498 standards at true equilibrium. Importantly, the exercise also demonstrates that if setup-specific
499 exchange fractionation factors $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ are used, the final Δ¹⁷O values should remain
500 consistent across laboratories. We suspect that the exchange fractionation factors used in the
501 above cases were not set-up specific causing the inter-laboratory discrepancies. Further

502 detailed analysis of the standards using the present approach will help validate and strengthen
503 this methodology.

504 **5. Conclusions**

505 We have experimentally verified that oxygen fractionation factors and exponent during Pt-
506 catalysed CO₂-O₂ exchange in reactors have a combined influence of oxygen isotope-exchange
507 between CO₂ and O₂ and thermal-gradient-induced kinetic fractionation. The fractionation
508 factors and exponent determined from uniformly heated samples show good agreement with
509 the theoretical values based on equilibrium modelling. This gives definitive evidence that
510 isotope exchange reaches to complete isotopic equilibrium during uniform heating whereas the
511 steady state achieved in exchange reactors with cold zones is shifted away from isotopic
512 equilibrium even though thermodynamic equilibrium states were attained. Further θ , supposed
513 to be an intrinsic property of the exchange process, is observed to be varying because of these
514 kinetic effects. These kinetic effects can be further verified by analysing clumped isotopes in
515 exchanged CO₂ and O₂, which can provide a constraint on their effective exchange
516 temperatures. The determined fractionation factors of exchange with varying reactor-
517 configurations showed a defined exponential correlation with the hot zone volumes of the
518 reactor. Thus, we propose a new framework to standardize $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ accounting
519 for the hot and cold zones of the exchange reactor for accurate triple oxygen measurements in
520 CO₂.

521

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529 levels of CO₂.

530

531 **Appendix A. Supplementary material:** Table S1-S6 and Figure S1 in this study can be found
532 in Supplementary Information. They contain theoretical predictions of fractionation factors and

533 results of the Pt-catalysed exchange conducted using isotopically different gas mixtures across
534 a range of temperatures and reactor geometries.

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558 **References**

- 559 Adnew, G.A., Workman, E., Janssen, C., Röckmann, T., 2022. Temperature dependence of isotopic
560 fractionation in the CO₂ -O₂ isotope exchange reaction. *Rapid Comm Mass Spectrometry* 36,
561 e9301. <https://doi.org/10.1002/rcm.9301>
- 562 Aron, P.G., Levin, N.E., Beverly, E.J., Huth, T.E., Passey, B.H., Pelletier, E.M., Poulsen, C.J., Winkelstern,
563 I.Z., Yarian, D.A., 2021. Triple oxygen isotopes in the water cycle. *Chemical Geology* 565,
564 120026. <https://doi.org/10.1016/j.chemgeo.2020.120026>
- 565 Bao, H., Cao, X., Hayles, J.A., 2016. Triple Oxygen Isotopes: Fundamental Relationships and
566 Applications. *Annu. Rev. Earth Planet. Sci.* 44, 463–492. [https://doi.org/10.1146/annurev-](https://doi.org/10.1146/annurev-earth-060115-012340)
567 [earth-060115-012340](https://doi.org/10.1146/annurev-earth-060115-012340)
- 568 Barkan, E., Affek, H.P., Luz, B., Bergel, S.J., Voarintsoa, N.R.G., Musan, I., 2019. Calibration of $\delta^{17}\text{O}$
569 and $^{17}\text{O}_{\text{excess}}$ values of three international standards: IAEA-603, NBS19 and NBS18. *Rapid*
570 *Comm Mass Spectrometry* 33, 737–740. <https://doi.org/10.1002/rcm.8391>
- 571 Barkan, E., Musan, I., Luz, B., 2015. High-precision measurements of $\delta^{17}\text{O}$ and $^{17}\text{O}_{\text{excess}}$ of NBS19 and
572 NBS18: $\delta^{17}\text{O}$ and $^{17}\text{O}_{\text{excess}}$ values of NBS19 and NBS18. *Rapid Commun. Mass Spectrom.* 29,
573 2219–2224. <https://doi.org/10.1002/rcm.7378>
- 574 Barkan, E., Luz, B., 2012. High-precision measurements of $^{17}\text{O}/^{16}\text{O}$ and $^{18}\text{O}/^{16}\text{O}$ ratios in CO₂. *Rapid*
575 *Comm Mass Spectrometry* 26, 2733–2738. <https://doi.org/10.1002/rcm.6400>
- 576 Bergel, S.J., Barkan, E., Stein, M., Affek, H.P., 2020. Carbonate $^{17}\text{O}_{\text{excess}}$ as a paleo-hydrology proxy:
577 Triple oxygen isotope fractionation between H₂O and biogenic aragonite, derived from
578 freshwater mollusks. *Geochimica et Cosmochimica Acta* 275, 36–47.
579 <https://doi.org/10.1016/j.gca.2020.02.005>
- 580 Bhattacharya, S.K., Thiemens, M.H., 1989. Effect of Isotopic Exchange upon Symmetry Dependent
581 Fractionation in the O + CO C O 2 Reaction. *Zeitschrift Für Naturforschung A*, 44, 811–813.
- 582 Bigeleisen, J., Mayer, M.G., 1947. Calculation of Equilibrium Constants for Isotopic Exchange
583 Reactions. *The Journal of Chemical Physics* 15, 261–267. <https://doi.org/10.1063/1.1746492>
- 584 Fosu, B.R., Subba, R., Peethambaran, R., Bhattacharya, S.K., Ghosh, P., 2020. Technical Note:
585 Developments and Applications in Triple Oxygen Isotope Analysis of Carbonates. *ACS Earth*
586 *Space Chem.* 4, 702–710. <https://doi.org/10.1021/acsearthspacechem.9b00330>
- 587 Furtenbacher, T., Harvey, A.H., Császár, A.G., 2025. Improved Partition Functions and Related
588 Thermochemical Quantities for the 16O2 and H216O Molecules. *Journal of Physical and*
589 *Chemical Reference Data* 54, 033103. <https://doi.org/10.1063/5.0273661>
- 590 Hofmann, M.E.G., Pack, A., 2010. Technique for High-Precision Analysis of Triple Oxygen Isotope
591 Ratios in Carbon Dioxide. *Anal. Chem.* 82, 4357–4361. <https://doi.org/10.1021/ac902731m>
- 592 Huang, X., Schwenke, D.W., Freedman, R.S., Lee, T.J., 2017. Ames-2016 line lists for 13 isotopologues
593 of CO₂: Updates, consistency, and remaining issues. *Journal of Quantitative Spectroscopy*
594 *and Radiative Transfer* 203, 224–241. <https://doi.org/10.1016/j.jqsrt.2017.04.026>
- 595 Huth, T.E., Passey, B.H., Cole, J.E., Lachniet, M.S., McGee, D., Denniston, R.F., Truebe, S., Levin, N.E.,
596 2022. A framework for triple oxygen isotopes in speleothem paleoclimatology. *Geochimica et*
597 *Cosmochimica Acta* 319, 191–219. <https://doi.org/10.1016/j.gca.2021.11.002>
- 598 Irikura, K.K., 2007. Experimental Vibrational Zero-Point Energies: Diatomic Molecules. *Journal of*
599 *Physical and Chemical Reference Data* 36, 389–397. <https://doi.org/10.1063/1.2436891>
- 600 Janssen, C., Tuzson, B., 2010. Isotope Evidence for Ozone Formation on Surfaces. *J. Phys. Chem. A*
601 114, 9709–9719. <https://doi.org/10.1021/jp1017899>
- 602 Jones, R.C., Furry, W.H., 1946. The Separation of Isotopes by Thermal Diffusion. *Rev. Mod. Phys.* 18,
603 151–224. <https://doi.org/10.1103/revmodphys.18.151>
- 604 Katakis, D., Taube, H., 1962. Some Photochemical Reactions of O₃ in the Gas Phase. *The Journal of*
605 *Chemical Physics* 36, 416–422. <https://doi.org/10.1063/1.1732521>
- 606 Koren, G., Schneider, L., Van Der Velde, I.R., Van Schaik, E., Gromov, S.S., Adnew, G.A., Mrozek
607 Martino, D.J., Hofmann, M.E.G., Liang, M., Mahata, S., Bergamaschi, P., Van Der Laan-Luijkx,

608 I.T., Krol, M.C., Röckmann, T., Peters, W., 2019. Global 3-D Simulations of the Triple Oxygen
609 Isotope Signature $\Delta^{17}\text{O}$ in Atmospheric CO_2 . *JGR Atmospheres* 124, 8808–8836.
610 <https://doi.org/10.1029/2019JD030387>

611 Laskar, A.H., Mahata, S., Bhattacharya, S.K., Liang, M., 2019. Triple Oxygen and Clumped Isotope
612 Compositions of CO_2 in the Middle Troposphere. *Earth and Space Science* 6, 1205–1219.
613 <https://doi.org/10.1029/2019EA000573>

614 Laskar, A.H., Mahata, S., Liang, M.-C., 2016. Identification of Anthropogenic CO_2 Using Triple Oxygen
615 and Clumped Isotopes. *Environ. Sci. Technol.* 50, 11806–11814.
616 <https://doi.org/10.1021/acs.est.6b02989>

617 Laskar, A.H., Maurya, A.S., Singh, V., Gurjar, B.R., Liang, M.-C., 2020. A new perspective of probing the
618 level of pollution in the megacity Delhi affected by crop residue burning using the triple
619 oxygen isotope technique in atmospheric CO_2 . *Environmental Pollution* 263, 114542.
620 <https://doi.org/10.1016/j.envpol.2020.114542>

621 Liang, M.-C., Laskar, A.H., Barkan, E., Newman, S., Thiemens, M.H., Rangarajan, R., 2023. New
622 constraints of terrestrial and oceanic global gross primary productions from the triple oxygen
623 isotopic composition of atmospheric CO_2 and O_2 . *Sci Rep* 13, 2162.
624 <https://doi.org/10.1038/s41598-023-29389-z>

625 Liang, M.-C., Mahata, S., Laskar, A.H., Bhattacharya, S.K., 2017a. Spatiotemporal Variability of Oxygen
626 Isotope Anomaly in near Surface Air CO_2 over Urban, Semi-Urban and Ocean Areas in and
627 around Taiwan. *Aerosol Air Qual. Res.* 17, 706–720.
628 <https://doi.org/10.4209/aaqr.2016.04.0171>

629 Liang, M.-C., Mahata, S., Laskar, A.H., Thiemens, M.H., Newman, S., 2017b. Oxygen isotope anomaly
630 in tropospheric CO_2 and implications for CO_2 residence time in the atmosphere and gross
631 primary productivity. *Sci Rep* 7, 13180. <https://doi.org/10.1038/s41598-017-12774-w>

632 Ma, Y., Peng, L., Zhang, H., Yu, J.-G., 2014. The potential energy surfaces of the ground and excited
633 states of carbon dioxide molecule. *Russ. J. Phys. Chem.* 88, 2339–2347.
634 <https://doi.org/10.1134/S0036024414130287>

635 Mahata, S., Bhattacharya, S.K., Wang, C.-H., Liang, M.-C., 2013. Oxygen Isotope Exchange between O_2
636 and CO_2 over Hot Platinum: An Innovative Technique for Measuring $\Delta^{17}\text{O}$ in CO_2 . *Anal. Chem.*
637 85, 6894–6901. <https://doi.org/10.1021/ac4011777>

638 Miller, M.F., Pack, A., 2021a. Why Measure 17O ? Historical Perspective, Triple-Isotope Systematics
639 and Selected Applications. *Reviews in Mineralogy and Geochemistry* 86, 1–34.
640 <https://doi.org/10.2138/rmg.2021.86.01>

641 Miller, M.F., Pack, A., 2021b. Why Measure 17O ? Historical Perspective, Triple-Isotope Systematics
642 and Selected Applications. *Reviews in Mineralogy and Geochemistry* 86, 1–34.
643 <https://doi.org/10.2138/rmg.2021.86.0>

644 Passey, B.H., Ji, H., 2019. Triple oxygen isotope signatures of evaporation in lake waters and
645 carbonates: A case study from the western United States. *Earth and Planetary Science Letters*
646 518, 1–12. <https://doi.org/10.1016/j.epsl.2019.04.026>

647 Passey, B.H., Hu, H., Ji, H., Montanari, S., Li, S., Henkes, G.A., Levin, N.E., 2014. Triple oxygen isotopes
648 in biogenic and sedimentary carbonates. *Geochimica et Cosmochimica Acta* 141, 1–25.
649 <https://doi.org/10.1016/j.gca.2014.06.006>

650 Prokhorov, I., Kluge, T., Janssen, C., 2019. Optical clumped isotope thermometry of carbon dioxide.
651 *Sci Rep* 9, 4765. <https://doi.org/10.1038/s41598-019-40750-z>

652 Richet, P., Bottinga, Y., Javoy, M., 1977. A review of Hydrogen, Carbon, Nitrogen, Oxygen, Sulphur, and
653 Chlorine Stable Isotope Fractionation among gaseous molecules. *Annu. Rev. Earth Planet. Sci.*
654 5, 65–110.

655 Sha, L., Dang, H., Wang, Y., Wassenburg, J.A., Baker, J.L., Li, H., Sinha, A., Ait Brahim, Y., Wu, N., Lu, Z.,
656 Yang, C., Dong, X., Lu, J., Zhang, H., Mahata, S., Cai, Y., Jian, Z., Cheng, H., 2024a. Triple
657 oxygen isotope reveals insolation-forced tropical moisture cycles. *Sci. Adv.* 10, eadp7855.
658 <https://doi.org/10.1126/sciadv.adp7855>

659 Sha, L., Dang, H., Wang, Y., Wassenburg, J.A., Baker, J.L., Li, H., Sinha, A., Ait Brahim, Y., Wu, N., Lu, Z.,
660 Yang, C., Dong, X., Lu, J., Zhang, H., Mahata, S., Cai, Y., Jian, Z., Cheng, H., 2024b. Triple
661 oxygen isotope reveals insolation-forced tropical moisture cycles. *Sci. Adv.* 10, eadp7855.
662 <https://doi.org/10.1126/sciadv.adp7855>

663 Sha, L., Mahata, S., Duan, P., Luz, B., Zhang, P., Baker, J., Zong, B., Ning, Y., Brahim, Y.A., Zhang, H.,
664 Edwards, R.L., Cheng, H., 2020. A novel application of triple oxygen isotope ratios of
665 speleothems. *Geochimica et Cosmochimica Acta* 270, 360–378.
666 <https://doi.org/10.1016/j.gca.2019.12.003>

667 Sha, Lijuan, Wassenburg, J.A., Sha, Lifan, Li, Y., Zhou, S., Liang, Q., Zhao, J., Ruan, J., Li, H., Zhao, X.,
668 Duan, P., Wang, J., Lu, J., Zhang, H., Kathayat, G., Mahata, S., Ban, F., Li, T.-Y., Cheng, H., 2023.
669 Variations in triple oxygen isotope of speleothems from the Asian monsoon region reveal
670 moisture sources over the past 300 years. *Commun Earth Environ* 4, 384.
671 <https://doi.org/10.1038/s43247-023-01043-6>

672 Spielfiedel, A., Feautrier, N., Cossart-Magos, C., Chambaud, G., Rosmus, P., Werner, H.-J., Botschwina,
673 P., 1992. Bent valence excited states of CO₂. *The Journal of Chemical Physics* 97, 8382–8388.
674 <https://doi.org/10.1063/1.463408>

675 Tashkun, S.A., Harvey, A.H., 2025. Partition Functions and Ideal-Gas Thermodynamics for Carbon
676 Dioxide. *Journal of Physical and Chemical Reference Data* 54, 023102.
677 <https://doi.org/10.1063/5.0276615>

678 Thiemens, M.H., Chakraborty, S., Dominguez, G., 2012. The Physical Chemistry of Mass-Independent
679 Isotope Effects and Their Observation in Nature. *Annu. Rev. Phys. Chem.* 63, 155–177.
680 <https://doi.org/10.1146/annurev-physchem-032511-143657>

681 Thiemens, M.H., Chakraborty, S., Jackson, T.L., 2014. Decadal $\Delta^{17}\text{O}$ record of tropospheric CO₂ :
682 Verification of a stratospheric component in the troposphere. *JGR Atmospheres* 119, 6221–
683 6229. <https://doi.org/10.1002/2013JD020317>

684 Thiemens, M.H., Heidenreich, J.E., 1983. The Mass-Independent Fractionation of Oxygen: A Novel
685 Isotope Effect and Its Possible Cosmochemical Implications. *Science* 219, 1073–1075.
686 <https://doi.org/10.1126/science.219.4588.1073>

687 Urey, H.C., 1947. The thermodynamic properties of isotopic substances. *J. Chem. Soc.* 562.
688 <https://doi.org/10.1039/jr9470000562>

689 Wei, Y., Yan, H., Peng, Y., Han, S., Bao, H., 2024. Thermal-gradient-induced isotope fractionation
690 during CO₂-O₂ triple oxygen isotope exchange. *Geochimica et Cosmochimica Acta* 370, 29–
691 40. <https://doi.org/10.1016/j.gca.2024.02.010>

692 Wostbrock, J.A.G., Cano, E.J., Sharp, Z.D., 2020. An internally consistent triple oxygen isotope
693 calibration of standards for silicates, carbonates and air relative to VSMOW2 and SLAP2.
694 *Chemical Geology* 533, 119432. <https://doi.org/10.1016/j.chemgeo.2019.119432>

695 Yu, S., Drouin, B.J., Miller, C.E., 2014. High resolution spectral analysis of oxygen. IV. Energy levels,
696 partition sums, band constants, RKR potentials, Franck-Condon factors involving the
697 $X^3\Sigma^-X^3\Sigma^-$, $a^1\Delta_g$ and $b^1\Sigma^+b^1\Sigma^+$ states. *The Journal of Chemical Physics* 141, 174302.
698 <https://doi.org/10.1063/1.4900510>

699 Yu, S., Miller, C.E., Drouin, B.J., Müller, H.S.P., 2012. High-resolution spectral analysis of oxygen. I.
700 Isotopically invariant Dunham fit for the $X^3\Sigma^-_g$, $a^1\Delta_g$, and $b^1\Sigma^+_g$ states. *The Journal*
701 *of Chemical Physics* 137, 024304. <https://doi.org/10.1063/1.4719170>

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707 **Figures**

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710 Figure 1. Schematic showing the in-house vacuum setup for Pt-catalysed CO₂-O₂ exchange for
711 triple oxygen measurements in CO₂

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714 Figure 2. (a) $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and (b) $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values determined with varying cold zone volumes
715 viz. 78.9%, 62.2%, 45% and 0%. The deviation from the theoretical predictions (presented in
716 Supplementary Information Table S1) increases with increasing cold zones. The shaded regions
717 along the theoretical curves represent the uncertainties.

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720 Figure 3. The oxygen isotope fractionation factors, $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ (dashed lines) and $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$
721 (solid lines), determined in four different experimental set-ups with varying cold zone volumes
722 within 600 to 1000 °C temperature range. The crossover temperatures, marked by vertical
723 dotted lines, increase as the cold zone volume decreases from 78.9% to 62.2% to 45%. In the
724 sealed quartz tubes with uniform heating, the $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values are closely aligned with the
725 theoretical values and do not show any crossing over of $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$.

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730 Figure 4. The $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ determined with varying cold zone volumes viz. 78.9%,
731 62.2%, 45% and 0% within 600 to 1000°C temperature range. When no cold zone is present,
732 the $\theta_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values are closely aligned with the theoretical value of 0.529. The curves show the
733 relationship predicted by Wei et al., 2024 using our revised equilibrium $^{18}\alpha_{\text{CO}_2/\text{O}_2}$ value and
734 $\theta_{\text{O}_2}^{\text{TG}}$ from their diffusion experiments viz., 0.48 (grey solid line), 0.5 (black solid line) and 0.52
735 (grey dotted line).

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739 Figure 5. Exponential relations of (a) $^{18}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ and (b) $^{17}\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ with varying hot zone volume
740 fraction at 750 °C. Selected $\alpha_{\text{CO}_2/\text{O}_2}^{\text{Pt}}$ values from previous studies with similar to our
741 experimental setups are also shown, assuming a 50:50 ratio between the heated and cold
742 sections of the reactors. The data from Adnew et al., 2022, in which the authors stated
743 exclusively that approximately half of the reactor was within the heating zone, are in good
744 agreement with our fitted line. For the other studies, hot and cold zone proportions need to be
745 constrained for the reactor geometry. The goodness of fit was evaluated using weighted
746 coefficient of determination (R_w^2).