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1 **Kaolin as a potential lithium source: a preliminary study of the**
2 **Torniella ceramic raw material deposit (southern Tuscany)**


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17 ABSTRACT

18 The Torniella ceramic raw material deposit (southern Tuscany) is hosted within highly altered rhyolites of the Tertiary Tuscan magmatic province.
19 Locally, kaolinite-rich veins occur, consisting of >80% kaolinite, with minor quartz and traces of svanbergite, $\text{SrAl}_3(\text{PO}_4)(\text{SO}_4)(\text{OH})_6$. The bulk vein
20 material shows interesting lithium (Li) contents, ranging from 1600 to 3700 mg/kg, with local enrichment up to 6000 mg/kg. Preliminary experiments
21 to assess the potential for Li extraction resulted in negligible release (0.2-0.4%) at room temperature (reaction with pure water, 1M MgCl_2 solution,
22 and 1M HCl solution). Roasting at 600 °C and leaching with oxalic acid resulted in up to 60% Li recovery. Conversely, roasting at 850 °C with sodium
23 sulfate and water leaching yielded a lower Li recovery (~40%). These results suggest that Li can be effectively mobilised only after thermal activation.
24 While the local resource is limited and the deposit itself is unlikely to be economically viable, this study acknowledges the potential of Li recovery as
25 a by-product of kaolin extraction for the ceramic industry.

26
27 **KEYWORDS:** lithium, kaolin, Torniella, lithium extraction.

28

29 INTRODUCTION

30 Lithium (Li) is included among the European Union list of critical raw materials primarily due to its significance in battery
31 technologies ([European Parliament, 2024](#)). At a global scale, the main producers of Li ores are Argentina, Australia, Chile,
32 China, and Zimbabwe ([U.S. Geological Survey, 2026](#)). Conventionally, Li extraction has relied on two primary resources:
33 continental brines (*salares*) and granitic pegmatites ([Benson et al., 2025](#)). A third type of deposits, referred to as either clay-
34 type deposits ([Bowell et al., 2020](#)) or volcano-sedimentary deposits ([Putzolu et al., 2025](#)), has recently attracted increasing
35 attention, especially because some prospects, such as Thacker Pass (Nevada), or Jadar (Serbia), contain estimated reserves
36 exceeding 1 Mt Li. Jadar represents a notable exception, where the Li-bearing phase is the unique mineral jadarite. However,
37 in the majority of clay-type deposits, Li is predominantly associated with clay minerals or micas, where it frequently substitutes
38 for Mg^{2+} in octahedral sites, as observed in hectorite ($\text{Na}_{0.3}(\text{Mg},\text{Li})_3\text{Si}_4\text{O}_{10}(\text{OH})$), or tainiolite ($\text{KLiMg}_2\text{Si}_4\text{O}_{10}\text{F}$). Finally, lithium
39 has been observed to occur adsorbed on the clay mineral surfaces or as an interlayer cation (e.g., [Benson et al., 2023](#); [Yan et](#)
40 [al., 2025](#)). In principle, kaolinite is not considered to be a favourable host for lithium (e.g., [Starkey, 1982](#)). Substitution in the
41 octahedral sites is limited by the large charge difference between Li^+ and Al^{3+} , while the theoretical absence of interlayer
42 cations and the relatively low cation exchange capacity should restrict Li incorporation through adsorption. Indeed, sorption
43 experiments by Hoyer et al. ([2015](#)) found that, under identical conditions, kaolinite takes up about one third of lithium with
44 respect to bentonite and zeolite. Nonetheless, natural kaolinite containing hundreds to even thousands of $\mu\text{g/g}$ Li has been
45 reported ([Hoyer et al., 2015](#); [Hao et al., 2025](#)), and there has been some recent attention, especially from Chinese researchers,
46 to the potential recovery of Li from kaolinite-bearing rocks (e.g., [Zhong et al., 2024](#)).

47 A joint research group at the Department of Earth Sciences, University of Florence, and at the Institute of Geosciences and
48 Georesources (CNR) was recently involved in two companion projects (ARTÙ and LiCycle) aimed at the recovery of lithium
49 from mine by-products and industrial non-compliant materials generated along the ceramic industry supply chain. Specifically,
50 one of the goals of the ARTÙ project is exploring the potential recovery of lithium and other critical metals from rocks quarried

51 at the ceramic raw material mines of *La Crocetta* (Porto Azzurro, LI, Elba Island) and *I Piloni* (Torniella, GR, Tuscany), run by
52 the EURIT Srl company (<http://euritonline.com/prodotti.html>). Run-of-the-mine bulk rocks, including rhyolites with varying
53 degrees of alteration from Torniella mine and aplites (commercially known as eurite) from Elba Island, were characterised from
54 petrographic, mineralogical, and geochemical perspectives. Preliminary studies on these lithotypes evidenced that no specific
55 mineralogical phases associated with lithium could be identified, while optical emission spectrometry (ICP-OES) detected Li
56 contents ranging from 35 to 180 mg/kg in rhyolites from Torniella, and between 119 and 169 mg/kg in aplites from Elba Island
57 ([Finto, 2025](#)). These values exceed the mean Li content of the upper continental crust (41 ± 6 mg/kg, [Hu & Gao, 2008](#)), and
58 are consistent with the documented Li anomaly within the products of the Tuscan magmatic province ([Dini et al., 2022](#)).
59 Notably, elevated Li contents were detected in samples from apparently pure kaolin veins at Torniella, up to 3000 mg/kg ([Finto,](#)
60 [2025](#)), i.e. markedly higher than the known Li anomaly within products of the Tuscan magmatic province ([Dini et al., 2022](#)). In
61 this study, further laboratory investigations were conducted on pure kaolin veins from Torniella mine to assess lithium
62 concentrations in these secondary assemblages and to explore the application of sustainable Li extraction methods. It should
63 be noted that these results are preliminary; further analyses are required for a comprehensive characterisation.

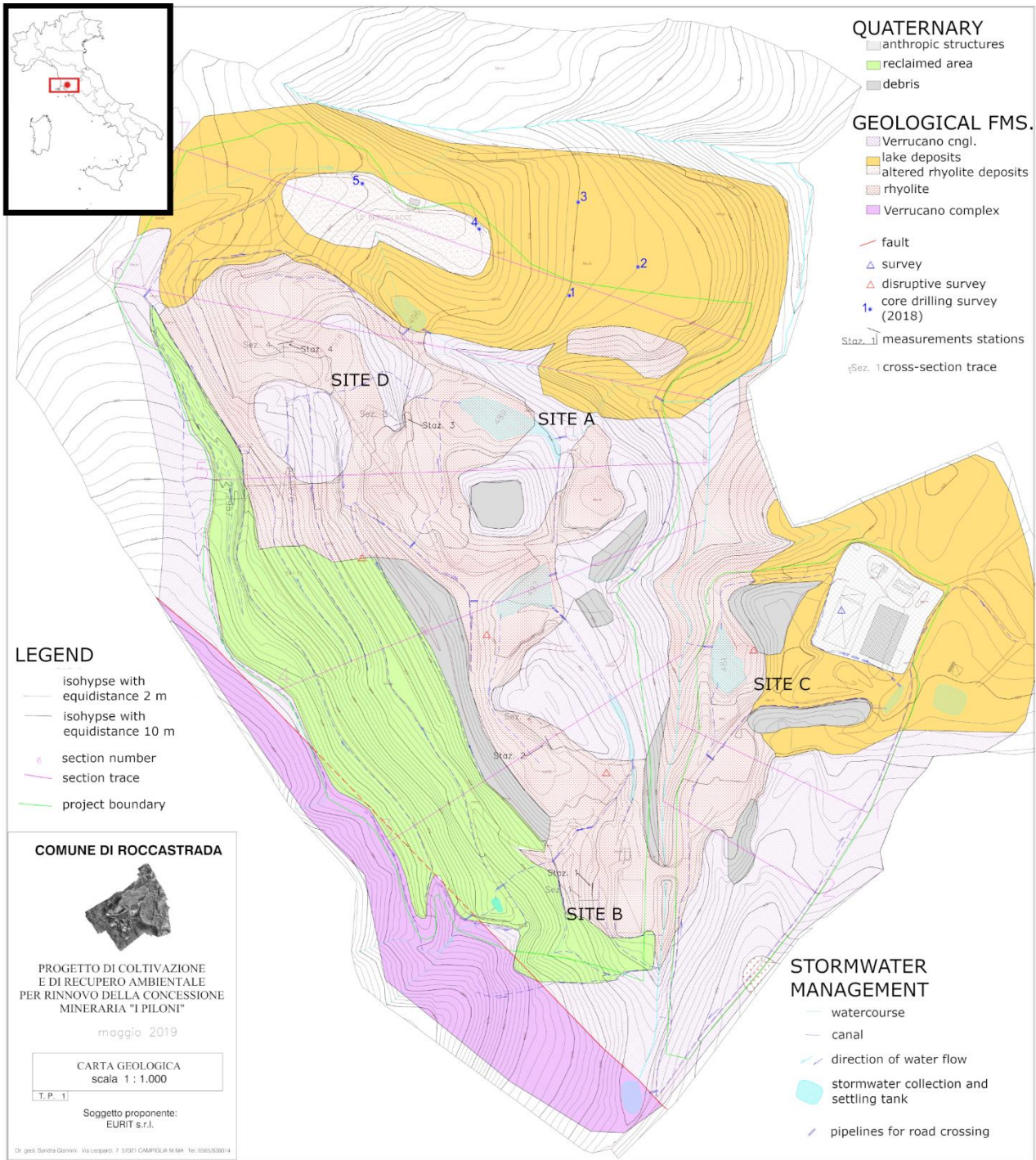
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65 **GEOLOGICAL CONTEXT**

66 The Torniella (*I Piloni*) mine near Roccastrada (GR) is one of the most significant deposits of kaolin in Italy ([Lombardi & Mattias,](#)
67 [1987](#); [Viti et al., 2007](#); [Sammuri & Scapigliati, 2022](#)) (Fig. 1). It is developed within rhyolitic rocks of the Tuscan Magmatic
68 Province, which are interpreted as the product of crustal anatexis and represent one of the most evolved and silica-rich
69 magmatic products of the province ([Mazzuoli, 1967](#); [Pinarelli et al., 1989](#)). The volcanic rocks crop out south of Torniella as a
70 NW-SE elongated body, roughly 3 km wide, whose western margin is controlled by a tectonic contact associated with a NW-SE
71 fault system; later SW-NE faults intersect the volcanic body, locally influencing the structural framework of the area. The
72 rhyolites are affected by variable degrees of hydrothermal alteration, producing lithologies that range from relatively well-
73 preserved volcanic rocks to strongly kaolinized materials. At present, a systematic description of the spatial distribution of
74 alteration is lacking. Broadly speaking, the phenomenon is associated with fractures mainly controlled by the Apennine and
75 anti-Apennine fault system ([Mazzuoli, 1967](#); [Bertolani & Loschi Ghittoni, 1989](#)). The resulting mineral assemblage is relatively
76 uniform, characterised by secondary kaolinite and alunite (and other minerals of the alunite supergroup), along with relict
77 primary minerals such as quartz and sanidine. Alunite has been mined in the past, but nowadays it compromises the quality of
78 the ore. Based on its isotopic composition ($\delta^{34}\text{S}$ values ranging from +6.7 to +9.6 per mil), [Cortecci et al. \(1981\)](#) suggested
79 that alunite at Torniella results from supergene oxidation of sulfides. However, there is no evidence that sulfides were ever
80 present at Torniella. An alternative explanation is the oxidation of H_2S upon reaction of ascending H_2S -bearing steam with cool
81 groundwater (steam-heated environment; cf. [Rye et al., 1992](#); [Hedenquist & Arribas, 2022](#)). The kaolinite-alunite acid-sulfate
82 alteration assemblage is comparable to that described in the Tolfa district ([Lombardi & Mattias, 1979](#)). The uneven distribution
83 of alunite at the Torniella mine reflects local variations in fluid chemistry, particularly in the availability of sulfur components

84 ([Lattanzi et al., 2001](#)). In addition to the primary hydrothermal deposition, minor occurrences of sedimentary kaolinite have
85 been identified within small, post-orogenic lacustrine basins to the south of the mine ([Bertolani & Loschi Ghittoni, 1989](#)). These
86 deposits were formed through the erosion and redeposition of kaolin derived from altered rhyolitic rocks. Alunite is scarce or
87 absent in these occurrences, which improves their economic quality ([Lombardi & Mattias, 1987](#); [Bertolani & Loschi Ghittoni,](#)
88 [1989](#)).

89 Currently, Torniella mine is exploited in different sectors, characterised by rhyolites from intact to intensely altered.
90 Specifically, sectors B and D of the mine exhibit considerable lithological variability, ranging from well-preserved rhyolites to
91 strongly kaolinized materials. In sector B, the rhyolitic formation exhibits localised fracture systems containing kaolin veins.
92 The fractures are predominantly sub-vertical, although sub-horizontal fractures with variable orientations are also present.
93 Kaolin veins range in thickness from a few millimetres to 2-3 cm, with isolated others reaching 4-5 cm (Fig. 2). They are
94 currently considered part of the mineable material. The relatively low iron content (see Tab. 3 below) is advantageous for
95 industrial applications, whereas the local occurrence of sulfate minerals is a disadvantage. Petrographic observations indicate
96 that the Torniella rhyolites are characterised by a porphyritic texture with a fractured glassy groundmass hosting abundant
97 phenocrysts of quartz and K-feldspar (up to centimetric), with minor occurrence of plagioclase and biotite flakes ([Viti et al.,](#)
98 [2007](#); [Finto, 2025](#)). In the less altered samples, the primary mineral assemblage is largely preserved, although plagioclase and
99 K-feldspar locally exhibit early alteration features, including fracturing and partial replacement by secondary phases. In
100 strongly altered rhyolites, plagioclase and biotite are completely replaced, and the groundmass develops a microgranular
101 texture dominated by kaolinite, whereas quartz and K-feldspar phenocrysts remain as relic grains. In the fracture-hosted kaolin
102 veins, the mineral assemblage is dominated by kaolinite, with minor quartz and K-feldspar.



103

104

Fig. 1. Geological map of Torniella (I Piloni) mine. Courtesy of EURIT Srl company.



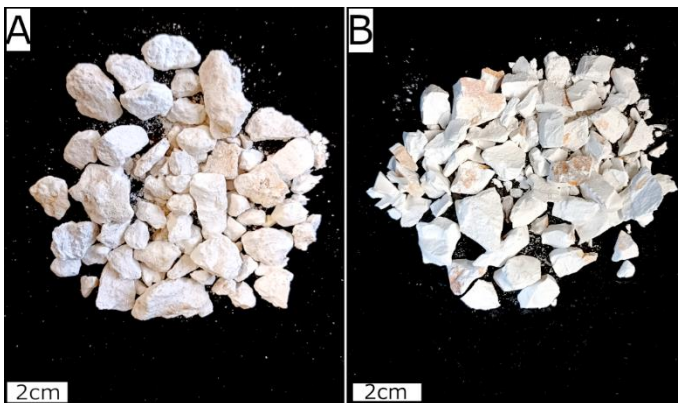
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106 Fig. 2. Kaolin veins, sector B, Torniella (I Piloni) mine.

107

108 MATERIALS & METHODS

109 Kaolin sampled from sector B veins at Torniella (about 2 kg in total) was collected along a fracture front approximately 12
 110 metres long and 10 metres high (scale assessed on the outcrop). The bulk colour of the sample is white, with occasional reddish
 111 streaks. Upon handling, it appeared consisting of a bulky, compact component, and a more loose, friable component. The two
 112 components were hand separated into two subsamples, KAO_013, the friable part, and KAO_013_BIS, the more compact part
 113 (Fig. 3). Both samples were dried in a temperature-controlled oven at 105 °C for 24h to remove moisture and were then ground
 114 to fine powders (~ 0.070 mm) for subsequent analyses.



115

116 Fig. 3. Kaolin sample from sector B, Torniella (I Piloni) mine: (a) KAO_013 (friable fraction), (b) KAO_013_BIS (compact
 117 fraction).

118

119 Microtextural, chemical and mineralogical characterization of kaolin

120 The microtextural and microchemical features of kaolin veins were studied using a Zeiss EVO MA15 scanning electron
 121 microscope coupled with an Oxford Aztec 5.0 EDS-EBSD spectrometer (Oxford ULTIM MAX 40 mm² EDS and Oxford Symmetry

122 EBSD) and Oxford INCA250 software. Working conditions were set as follows: accelerating voltage 15 kV, beam current 10 nA,
123 working distance 8.5 mm.

124 The bulk mineralogical composition of the two samples was assessed by X-ray powder diffraction (XRPD). The analysis was
125 conducted using a XRDynamic 500 diffractometer (Anton Paar), with Cu-K α radiation ($\lambda = 1.5406 \text{ \AA}$, 40 kV, 40 mA) over the
126 3-90°/2 θ range, step size 0.02°/2 θ , time per step 60 s. Phase identification was performed using DIFFRAC.EVA (Bruker)
127 software, with the Crystallography Open Database. Semi-quantitative analysis was carried out using DIFFRAC.TOPAS software
128 (Coelho, 2018). Anisotropic broadening is significant in kaolinite (Bish et al., 1989) and was modelled by applying spherical
129 harmonics as implemented in Topas 6 (Dinnebier et al., 2018). Furthermore, XRPD data were also employed to evaluate the
130 structural order-disorder of kaolinite. This parameter is relevant for this study because a higher crystallinity corresponds to a
131 well-ordered lattice with fewer structural defects and stronger cation bonding, which makes cationic exchange increasingly
132 challenging (Pineau et al., 2022). Finally, the mineral phase transformations of the bulk samples upon heating were studied
133 using Thermo-X-Ray Diffraction analysis with a XRDynamic 500 diffractometer (Anton Paar) and a Cu-K α radiation, ranging 3-
134 70°/2 θ , step size of 0.02°, time per step 60 s. The diffractometer was equipped with an Anton Paar HTK-1200N furnace for in
135 situ heating; measurements were performed in air with a heating rate of 10 °C/min.

136 The bulk chemical characterisation of the samples was performed using X-ray fluorescence spectrometry (XRF). Glass beads
137 with a diameter of 30 mm were prepared in an electric beader from 1 g of sample mixed with 10 g of Li₂B₄O₇ flux. The analysis
138 was performed using a ZSX Primus II WD-XRF (Rigaku Industrial Corporation), with a Rh source (4kW, 60mA) in a vacuum
139 chamber. Quantitative analysis of major, minor and trace elements (except Li) was performed using calibration curves created
140 and validated through a large set of international geostandards. Loss on ignition (LOI) was determined by heating three
141 replicates of 0.5 g of each sample at 950 °C for 1.5 hours.

142 Li contents in the kaolin samples were determined using Inductively Coupled Plasma–Optical Emission Spectrometry (ICP-
143 OES) following a total digestion protocol. This procedure is designed to ensure a gradual and complete dissolution of all mineral
144 phases within the sample matrix. Four representative aliquots (~0.1 g each) of each sample were digested in Teflon vessels
145 on a temperature-controlled hotplate, following sequential heating and evaporation steps using a mixture of concentrated
146 acids (HNO₃, HF, HCl) until complete dissolution was achieved (Balaram & Subramanyam, 2022). For testing analytical
147 accuracy, digestion was also performed on an international standard reference material (JR1, Geological Survey of Japan).
148 Procedural blank samples were also run to evaluate possible contamination. The final solutions were diluted with MilliQ for
149 elemental analysis on a Thermo Fisher iCAP PRO XP ICP-OES instrument (Thermo Fisher Scientific Inc, Massachusetts, USA),
150 operating with RF power 1.15 kW, nebuliser pressure 5 bar, sample uptake rate 45 rpm, and exposure time 5 s. Measurements
151 were obtained in radial mode using the Li emission line at 670.78 nm. Measurement precision across sample replicates showed
152 a relative standard deviation consistently below 5%, with accuracy to the standard within $\pm 7\%$.

153

154 **Leaching tests for Li extraction**

155 Conventional lithium recovery from clays relies on high-temperature mobilisation of Li via exchange with H⁺ ions from strong
156 acids (e.g., [Zhao et al., 2023](#)), entailing significant environmental and energy impacts. The present study involved the
157 application of procedures aimed at reducing the environmental impact, using reagents with a low ecological footprint to
158 minimise waste toxicity. Specifically, a preliminary compliance test, a room temperature cation exchange test and two
159 calcination and leaching protocols were applied. A further room temperature exchange test using a strong acid was included
160 in order to provide a direct comparison with less sustainable conditions. It is important to note that these tests are of exploratory
161 nature and not intended to be exhaustive.

162 The preliminary compliance test was conducted in accordance with the European Standard UNI EN 12457-2 ([UNI, 2002](#)) for
163 the water leaching of granular waste materials and sludge, using distilled water as the leaching agent with a solid-to-liquid ratio
164 of 5.0 g / 50 mL, and stirring at 300 rpm for 24 h at 21 °C. The room temperature exchange test was carried out using a 1 M
165 MgCl₂ solution as the exchange agent, with a solid-to-liquid ratio of 5.0 g / 50 mL, and stirring at 300 rpm for 24 h at 21 °C.
166 The calcination and leaching protocol developed by [Zhong et al. \(2024\)](#) proposes the utilisation of sodium sulfate as an
167 alternative to conventional strong acids, such as H₂SO₄. Sodium sulfate is expected to have a reduced environmental impact
168 and the capacity to promote Li extraction whilst constraining the dissolution of aluminium and iron ions. The procedure was
169 originally applied to a material from the Yuxi mine in Yunnan province (China), containing 0.18 wt% Li₂O, which is said to be
170 mainly present in the kaolinite crystalline structure. A mixture of kaolin (45%) and Na₂SO₄ (55%) was calcined at 800 °C for
171 90 min with a heating rate of 10 °C/min. The calcined product was then leached in water at a solid-to-liquid ratio of 3.0 g / 15
172 mL. Leaching was performed at 90 °C for 60 min, with a stirring speed of 300 rpm. The leaching efficiency has been reported
173 to be 84%. The second calcination and leaching test was based on the procedure suggested by [Liu et al. \(2024\)](#), where strong
174 inorganic acids are replaced by a less impacting organic acid such as oxalic acid. The authors treated a complex material
175 composed of chlorite, kaolinite, diaspore, boehmite and muscovite, containing 0.75 wt% Li₂O, where the structural position of
176 lithium remains uncertain. The material was calcined without additives at 600 °C for 60 min (heating rate 10 °C/min) and
177 subsequently leached using a 1.2 M oxalic acid solution at a solid-to-liquid ratio of 1:8. The leaching process was conducted
178 at 80 °C for 180 min, with a stirring speed of 300 rpm. The leaching efficiency reported by [Liu et al.](#) was 91%. Finally, an
179 exchange test with a strong acid was conducted at room temperature, employing a 1 M HCl solution as the exchange agent,
180 with a solid-to-liquid ratio of 5.0 g / 50 mL, stirring at 300 rpm for 24 h at 21 °C.

181 Following centrifugation, Li concentration in each eluate was determined using ICP-OES under the same conditions as those
182 applied for the total lithium content. The recovery of Li was calculated using Equation 1,

$$183 \quad \varepsilon(Li, \%) = ((c_{Li \text{ leached}} * v_{leached}) / (c_{Li \text{ raw}} * m_{bulk \text{ raw}})) * 100 \quad (1)$$

184 Where ε represents the lithium recovery, $c_{Li \text{ leached}}$ represents the Li concentration in the eluate, $v_{leached}$ represents the
185 volume of the eluate, $c_{Li \text{ raw}}$ represents the lithium content in the raw material, $m_{bulk \text{ raw}}$ represents the mass of the sample
186 used in the leaching procedure.

187

188 **RESULTS**189 **Microtextural, chemical and mineralogical characterization of kaolin**

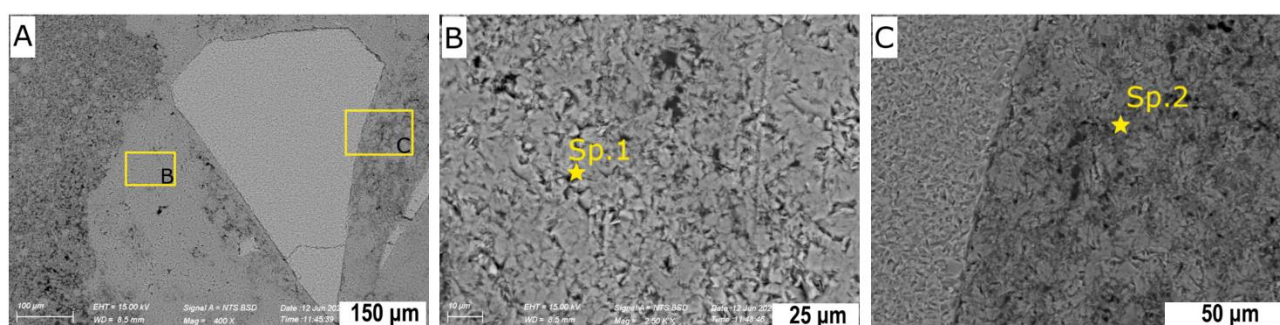
190 SEM imaging of thin sections of the kaolin veins revealed a microcrystalline texture (Fig. 4a). Semiquantitative EDS analyses
191 indicated the presence of Sr–Al sulfate minerals consistent with svanbergite ($\text{SrAl}_3(\text{PO}_4)(\text{SO}_4)(\text{OH})_6$), occurring both as
192 microcrystalline aggregates and dispersed within the kaolinitic matrix (Fig. 4c, Tab. 1). Locally, As-rich compositions
193 compatible with kemmlitzite ($\text{SrAl}_3(\text{AsO}_4)(\text{SO}_4)(\text{OH})_6$) were also detected.

194 Bulk XRD analysis of the kaolin-Li-rich samples indicates a composition dominated by kaolinite, with minor quartz and sodic
195 plagioclase in the friable fraction (KAO_013) (Fig. 5). Traces of svanbergite, $\text{SrAl}_3(\text{PO}_4)(\text{SO}_4)(\text{OH})_6$, a mineral of the alunite
196 supergroup, were detected in both samples (Tab. 2). No Li-bearing mineral phases were identified. The determination of the
197 kaolinite crystallinity index (KCI) from X-ray diffraction data is complicated by overlapping peaks from associated minerals and
198 amorphous phases. For example, the widely used Hinckley index (HI) is strongly affected by the presence of quartz, feldspar,
199 iron hydroxide gels, illite, smectite and halloysite (Aparicio & Galan, 1999). Consequently, the KCI could not be calculated for
200 KAO_013 due to the presence of interfering accessory phases, namely quartz. For KAO_013_BIS, the Aparicio-Galán-Ferrell
201 index (AGFI) was determined by fitting the peak heights of the 020, 1-10 and 11-1 reflections of kaolinite. The calculated AGFI
202 is 0.7, indicative of a highly defective kaolinite (Aparicio et al., 2006). For comparison, the Hinckley Index (HI), based on the
203 intensity of the same reflections, returned a value of 0.8, corresponding to a medium-defect kaolinite structure (Hinckley,
204 1962) (Tab. 2). Thermo-XRD analyses of the two kaolin samples show, as expected, that the characteristic 001 peak of
205 kaolinite (7.2 Å) starts decreasing its intensity above 500 °C due to dehydroxylation, and is completely missing at 600 °C,
206 possibly indicating the formation of metakaolinite (X-ray amorphous phase) (Fig. 6). At the same temperature, svanbergite is
207 no longer detected, as it underwent dehydration and loss of crystallinity.

208 XRF analyses of the two kaolin samples showed similar chemical compositions in terms of major and minor oxides, as well as
209 trace elements (Tab. 3, Tab. 4). The dominance of silica and alumina is fully consistent with a mineralogical composition
210 dominated by kaolinite. The relatively elevated Sr concentrations measured in both samples are attributable to the presence
211 of Sr-Al sulfates, as identified by XRPD.

212 The total Li content of the kaolin samples considered for the present explorative study was measured with ICP-OES and resulted
213 in 1667 mg/kg for KAO_013 and 3556 mg/kg for KAO_013_BIS (Tab. 4).

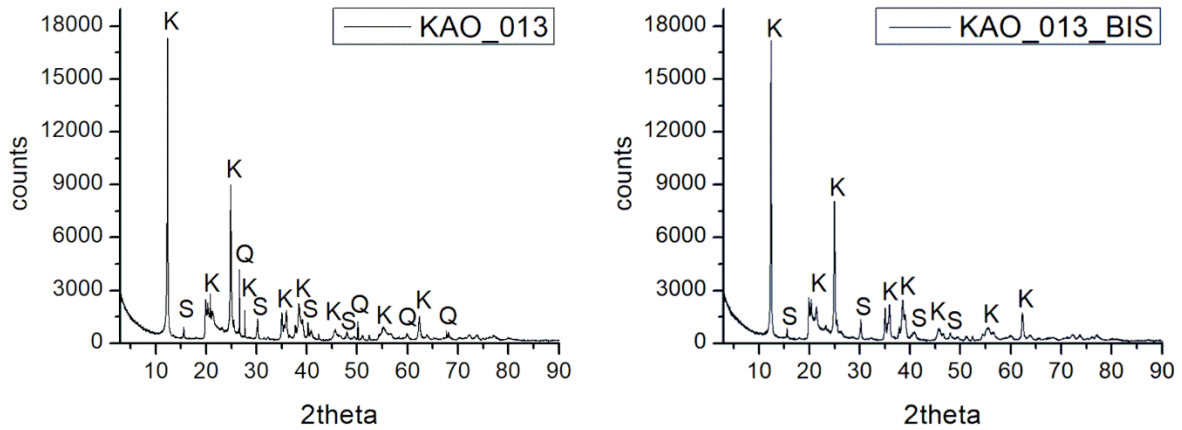
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215 Fig. 4. SEM secondary electron (SE) images of kaolin vein sample. Sp. 1 and Sp. 2 correspond to the points where EDS
 216 microanalyses were conducted. The semiquantitative composition (see Tab. 1) suggests for Sp.1 a Si-Al composition
 217 compatible with kaolinite; whereas Sp.2 suggests the presence of svanbergite (darker spots in Fig. 6c) in the kaolinitic matrix.

CHEMICAL COMPOSITION [wt%]	Al ₂ O ₃	SiO ₂	P ₂ O ₅	SO ₄	K ₂ O	SrO
KAO_003_Sp.1	46.7	53.3	-	-	-	-
KAO_003_Sp.2	37.6	15.6	11.4	16.0	0.6	15.3

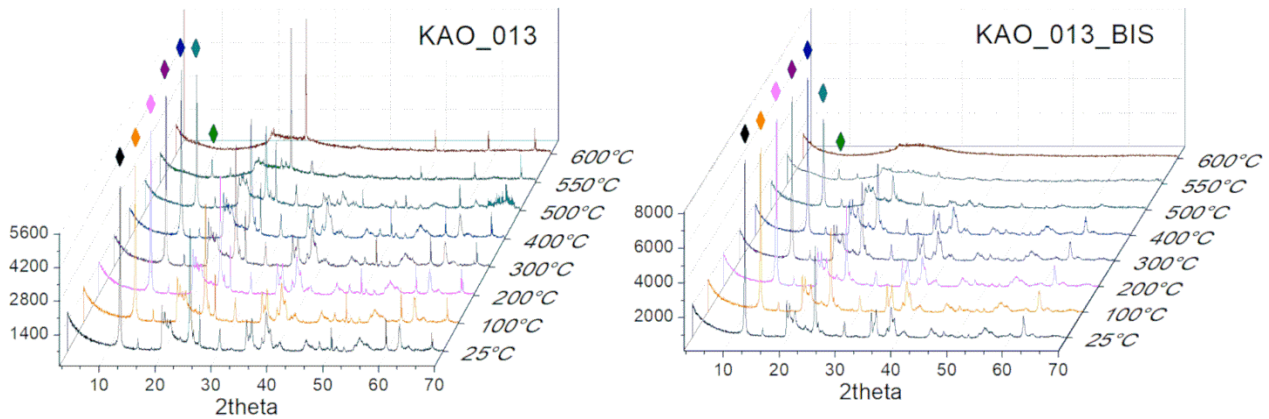
218 Tab. 1. Semiquantitative composition of kaolin vein sample obtained with SEM-EDS measurements.



219
 220 Fig. 5. XRD patterns for KAO_013 and KAO_013_BIS bulk raw powders; peak identification: K = kaolinite, S = svanbergite, Q
 221 = quartz.

	XRPD bulk semiquantitative analysis				kaolinite crystallinity index	
	clays (kaolinite)	alunite supergroup minerals	quartz	plagioclase	AGFI	HI
KAO_013	89%	3%	7%	<2%	N/A	N/A
KAO_013_BIS	97%	3%	<DL	<DL	0.7	0.8

222 Tab. 2. Bulk mineralogical composition of the kaolin samples (error 3-5%; LOD ~1%). Kaolinite crystallinity calculated with
 223 AGFI and HI indices.



224
 225 Fig.6. Thermo-XRD analysis results for KAO_013 and KAO_013_BIS, in the range of 25-600 °C. The kaolinite 001 peak is
 226 indicated for each diffractogram, if present.

ANALYTE	SiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	TiO ₂	P ₂ O ₅	SrO	LOI	TOT	SO ₃ *
UNIT	wt%	wt%	wt%	wt%	wt%	wt%	wt%	wt%	wt%	wt%	wt%	%	-	wt%
DL	0.50	0.80	0.03	2*10 ⁻³	0.03	0.02	0.30	0.02	0.01	0.02	0.5*10 ⁻³	-	-	-
KAO_013	43.88	38.46	0.39	0.01	<DL	0.07	<DL	0.14	0.11	0.53	0.97	14.66	99.21	0.75
KAO_013_BIS	40.14	41.65	0.17	0.01	<DL	0.07	<DL	0.08	0.03	0.57	1.06	15.10	98.87	0.82

227 Tab. 3. Major and minor oxide composition of the bulk kaolin samples. DL indicates the detection limit. * SO₃ contents
 228 correspond to the amount estimated to be present in svanbergite, calculated from the measured Sr concentrations; this
 229 contribution is included in the reported LOI values.

ANALYTE	V	Cr	Ni	Cu	Zn	Zr	Ba	Co	Rb	Y	Nb	Li *
UNIT	mg/kg	mg/kg	mg/kg	mg/kg	mg/kg	mg/kg	mg/kg	mg/kg	mg/kg	mg/kg	mg/kg	mg/kg
LOD	3	10	5	3	3	5	30	5	5	20	30	0.5*10 ⁻³
KAO_013	52	69	59	78	4	27	272	5	14	<DL	<DL	1712 ± 49
KAO_013_BIS	69	63	59	50	4	<DL	327	6	<DL	<DL	<DL	3628 ± 90

230 Tab. 4. Trace element contents in the bulk kaolin samples. DL indicates the detection limit. * Li content analysed by ICP-OES
 231 (emission line 670.78 nm), reported as the mean value between four different aliquots of the sample.

232 Leaching tests for Li extraction

233 The compliance test, the cation exchange test, and the room temperature leach with a 1M HCl yielded very low Li release.
 234 Specifically, the compliance test conducted in accordance with the European Standard UNI EN 12457-2 (room temperature
 235 batch reaction with pure water) showed a Li release of only 0.1% for both samples. The room temperature exchange test in a
 236 1M MgCl₂ solution achieved a recovery of Li of 0.2% for KAO_013 and KAO_013_BIS (Tab. 5). Finally, the room temperature
 237 exchange test in a 1M HCl solution achieved a recovery of Li of 0.4% for both samples (Tab. 5). Calcination of the kaolin
 238 samples at 800 °C with Na₂SO₄ and water leaching at 90 °C, yielded Li recoveries of 37% for KAO_013 and 22% for
 239 KAO_013_BIS (Tab. 5). XRPD analyses of the calcined powders revealed residual quartz and sodium sulfate phases in
 240 KAO_013, whereas only the sodium sulfate pattern was detected in KAO_013_BIS. This indicates that both kaolinite and
 241 svanbergite underwent complete transformation during the calcination process. However, Li extraction efficiencies remained
 242 comparatively low in relation to the findings reported by Zhong et al. (2024). In contrast, calcination at 600 °C without additives,
 243 followed by leaching in oxalic acid solution at 80 °C, resulted in significantly higher Li recovery rates, although these were still
 244 below the values reported by Liu et al. (2024): 58% for KAO_013 and 53% for KAO_013_BIS (Tab. 5). XRPD analyses of the
 245 calcined powders revealed residual quartz in KAO_013, whereas KAO_013_BIS was found to be completely amorphous.

	Room temperature exchange in MgCl ₂ solution	Calcination with Na ₂ SO ₄ + hot leaching in water solution	Calcination + hot leaching in oxalic acid solution	Room temperature exchange in HCl solution
KAO_013	0.2%	37%	58%	0.4%
KAO_013_BIS	0.2%	22%	53%	0.4%

246 Tab. 5. Li recovery (%) after exchange tests and leaching processes.

247

248 DISCUSSION

249 Kaolinite-rich veins from the Torniella rhyolite mine locally contain remarkably high Li contents, ranging from 1600 to 3700
250 mg/kg, with local enrichment of up to 6000 mg/kg. These concentrations are significantly higher than those observed in the
251 hosting rocks. Previous studies (summarised in [Lattanzi et al., 2001](#)) suggest a structurally controlled hydrothermal alteration
252 by ascending fluids, possibly originated in a steam-heated environment. There is no specific evidence about the Li-hosting
253 phase(s) in the unaltered rhyolite. The most probable candidates involved in Li mobilisation during hydrothermal alteration are
254 biotite and the glass phase (e.g. [Ellis et al., 2022](#)). Thus, vein minerals, and specifically kaolinite, may have concentrated the
255 Li mobilised from rhyolite alteration. Kaolinite has been identified as the most probable host phase for lithium within the veins,
256 although direct evidence of Li structural position is currently lacking. Li contents in Torniella kaolin veins are among the most
257 relevant documented for this mineral, and of the same order of magnitude of grades considered economic in some clay-type
258 Li deposits ([Bowell et al., 2020](#); [Xie et al., 2024](#)). Exchange tests at room temperature in pure water, MgCl₂ 1M and HCl 1M
259 resulted in low extraction efficiencies, approximately 0.2 - 0.4%. These results suggest that Li is not simply adsorbed onto the
260 kaolinite structure (e.g., surface defects), otherwise it would have reacted to the mild exchange conditions, being released in
261 the solution upon substitution with H⁺ or Mg²⁺. The thermally induced transformations of kaolinite during calcination appear to
262 play a key role in enhancing Li extraction efficiency. Between 400 °C and 600 °C, kaolinite undergoes dehydroxylation and
263 structural collapse of its layered sheets, leading to the formation of metakaolin. This behaviour was confirmed by thermo-XRD
264 analyses, which indicate that the transformation from kaolinite to X-ray amorphous phases in the Torniella samples is complete
265 at 600 °C. High-temperature transformations have been shown to enhance reactivity and Li mobility ([Dhar & Bishnoi, 2024](#);
266 [Liu et al., 2024](#)). Indeed, leaching tests following calcination yielded significant Li extraction, although considerably lower than
267 the >80% efficiencies documented in the reviewed literature. Despite the lower calcination temperature of 600 °C, the
268 combination of heating without additives followed by hot leaching in oxalic acid solution resulted in a higher Li recovery than
269 the calcination at 800 °C with Na₂SO₄, followed by hot water leaching. These results suggest that the exchange between H⁺
270 from the oxalic acid and Li⁺ is favoured in comparison to the exchange of lithium with Na⁺ from sodium sulfate. The
271 experimental procedures were conducted under conditions described as optimal in the extant literature, including reagent
272 concentrations, calcination temperature and duration, as well as leaching time and temperature ([Liu et al., 2024](#); [Zhong et al.,](#)
273 [2024](#)). The present study did not investigate whether alternative conditions, such as higher calcination temperatures or
274 extended reaction times, may yield enhanced Li extraction efficiencies for Torniella kaolin samples. Furthermore, although
275 kaolin represents the primary material in both the consulted literature and the present study, variations in accessory mineral
276 phases may influence the overall extraction behaviour. At this stage, given the exploratory nature of the study, the pursuit of
277 parameter optimisation was not deemed to be a priority. These results, although preliminary, suggest that Li-rich kaolinite may
278 represent a potential unconventional source of Li, particularly in the context of by-product recovery from kaolin extraction.

279

280 CONCLUSIONS

281 The rhyolite bodies currently exploited at the Torniella (*I Piloni*) mine, one of the most significant ceramic raw material deposits
282 in Italy, exhibit a moderate Li anomaly (up to 160 mg/kg), a characteristic that is commonly observed in the Tertiary Tuscan
283 magmatic province. Conversely, a sector of the mine is distinguished by the occurrence of centimetric veins, predominantly
284 composed of kaolinite, which exhibit significantly higher Li contents (1600 - 3700 mg/kg, with local enrichment up to 6000
285 mg/kg). Preliminary tests of Li recovery demonstrated a limited mobility of the metal. Room temperature reaction with water,
286 $MgCl_2$ and HCl, released approximately 0.2 - 0.4% Li, indicating that the metal is not weakly adsorbed, but structurally bound.
287 Thermal treatment and hot leaching yielded better results (22 - 58%), even if lower than those reported in the literature for the
288 same procedures. Specifically, the most positive results (up to ~60% recovery) were obtained with calcination at 600 °C and
289 hot leaching with oxalic acid. This technique has been shown to outperform calcination with Na_2SO_4 at 800 °C and water
290 leaching, suggesting that H^+ exchange in acidic conditions is more effective than Na^+ -based mechanisms. Additional
291 optimisation of process parameters, including temperature, duration, and reagent concentration, may yield better outcomes.
292 At the current state of knowledge, the lithium resource appears negligible, being limited to a small portion of the Torniella mine.
293 Furthermore, the presence of sulfur compounds (svanbergite, kemmlitzite) penalises the kaolin residue for ceramic
294 applications. Although not currently economically viable, the results of this study demonstrate a concrete potential for Li
295 recovery as a by-product from Li-rich kaolin. The final goal is to develop processes that preserve the properties of kaolin for
296 post-leaching industrial applications, paving the way for integrated resource recovery strategies. Future works would focus on
297 optimising extraction conditions and investigating the structural position of lithium within the kaolinitic matrix by means of
298 advanced analytical techniques.

299

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