1	Is the Earth Lazy? A review of work minimization in fault evolution
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- 11

11 Abstract

12 The principle of work minimization has been used in various forms to account for the 13 development of active fault systems within a wide range of tectonic settings. We review the 14 successes, challenges and implications learned from previous applications of work 15 minimization. Examination of the energy budget provides insight into the competing 16 influences of different processes within fault systems at a variety of scales. We present 17 each work budget component with considerations for numerical implementation. 18 Additionally, we demonstrate numerical implementation by solving for the energy budgets 19 and finding the most efficient propagation paths of a joint and several faults. Consideration 20 of the energy budget captures growth patterns that are consistent with both laboratory 21 observations and theoretical predictions using the energy release rate, G. Unlike using G, 22 work minimization is not limited to collinear growth. In comparison with predictions using 23 Coulomb failure planes, work minimization eliminates uncertainty in predicting fault 24 propagation by (1) evaluating the trade-off between tensile and shear failure ahead of the 25 fault and (2) avoiding the ambiguity introduced by the two potential Coulomb failure 26 planes. Application of work minimization to faulting demonstrates the effectiveness of this 27 approach in predicting both the orientation and timing of fault propagation.

28 1. Introduction

Crustal deformation results from complex plate motions acting on a variety of time scales, from the seconds required for earthquake ruptures to the millions of years required for mountain building. Brittle failure plays a large role in this deformation, as faults slip and grow to accommodate tectonic loading and dykes propagate due to internal magma pressure, for example. Predicting the growth of brittle structures is critical to unraveling many geophysical processes and is particularly important for understanding fault evolution.

Inglis (1913) provided fundamental insights into the role of concentrated stresses
around flaws, and their role in flaw growth and the subsequent failure of the material.
Using these insights, the growth of fractures can be predicted if the stress concentration
factor at the fracture tip meets the strength of the material. This concept is fundamental to

40 theories of linear elastic fracture mechanics and has been successful in predicting 41 conditions for and paths of opening mode propagation that is characteristic of joints, veins 42 and dikes in rock materials (e.g. Pollard and Aydin, 1988). Analysis of the propagation of 43 faults using this approach has been more problematic. The challenges arise from the 44 distributed nature of damage preceding fault growth and the subsequent linkage of cracks 45 to form fault surfaces. The deformational processes around a fault tip do not match the in-46 line growth pattern required for using stress concentration factors to predict the potential 47 for fault growth. In addition, while seismologists have been successful in using a local 48 energy balance at the fracture tip to model earthquake dynamics, strength heterogeneities 49 in the Earth make application of this approach over large areas and long time-scales 50 particularly difficult.

51 In contrast, work minimization provides a global approach to predicting failure by 52 postulating that the crust deforms in order to minimize the external tectonic work acting 53 on it. This deformation includes the propagation of faults and work minimization suggests 54 that fault propagation will occur in the direction that minimizes external energy. As a 55 consequence, a fault system may evolve over geologic time to efficiently accommodate 56 forcing from plate motions. A work minimization approach does not ignore fracture 57 mechanics, but rather drives mechanics based on a global energy budget, rather than solely 58 on a criterion local to the fault or fracture tip. Furthermore, fault growth along a 59 mechanically efficient path can still require that local conditions for failure are met at a 60 fault tip, though the timing and orientation of growth are based on the global energy 61 budget.

62 While the principle of work minimization has been applied in various forms in the 63 past, now is an ideal time to apply it to research on crustal deformation in general and fault behavior in particular. Current computing capabilities allow fault mechanics to be 64 65 considered and the full energy budget of deformation to be determined. This eliminates resurgent concerns regarding minimization approaches, such as that strain minimization 66 67 fails to capture the "complete energetics of the system" (Bird and Yuen, 1979). 68 Computational advances also allow for better constraints on and quantification of individual work budget components, including the energy required in work against friction, 69

70 energy required for propagation, energy converted into seismic waves, work against

71 gravity, and the internal strain energy of the deforming system.

72 In this review paper, we present the individual components of the energy budget 73 along with guides for numerical implementation of the calculations. Different approaches 74 that previous investigators have taken to predict fault evolution using the principle of work 75 minimization are outlined, as well as the historical debates about the appropriateness of 76 work minimization in the study of fault network evolution. To demonstrate the benefits of 77 analyzing the complete work budget, we apply work minimization to the growth of a joint 78 and a fault, and contrast their energy budgets. The paper concludes with discussion of the 79 insights provided by using a work minimization approach to investigate tectonic 80 deformation in general and fault system evolution in particular, as well as on-going 81 considerations for doing so and current challenges.

82 **2. Energy budget for fault evolution**

83 Prior to reviewing the contributions of previous investigations of fault evolution using 84 work minimization, we outline the individual components of the energy budget. Over the 85 years, various workers have used slightly different definitions of these energy terms and 86 utilized different subsets of the complete energy budget. Here we present a set of 87 consistent formulations for a complete energy budget associated with fault growth. This budget includes internal work of deformation around faults, Wint, work against gravity, 88 89 W_{grav} , work against friction along faults, W_{fric} , energy to create fault surfaces, W_{prop} , and 90 energy of ground shaking, *W*_{seis}. The sum of these terms equals the total work of the system, 91 which is the external work, W_{ext} , along the boundaries of the fault system (Fig. 1): 92 $W_{ext} = W_{int} + W_{grav} + W_{fric} + W_{prop} + W_{seis}$ Eq. 1 93 The internal work and work against gravity often are considered together as mechanical 94 work. Within fault systems, gravitational force plays a different role within contractional 95 and extensional systems (Dempsey et al., 2012), so separating these two work components 96 provides relevant insights. All work terms except for the energy consumed by creating 97 fault surfaces (W_{prop}) can be derived from the product of a force and a displacement within 98 the fault system. In addition to describing the theoretical underpinning of each component

99 of the work budget, we note special considerations for the numerical calculation of each100 component.

101 2.1 Internal work of deformation

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Internal work is the work of deformation within the fault system. *W_{int}* can be
measured as strain energy density, which is the product of stress and strain (e.g.
Timoshenko and Goodier, 1951; Jaeger et al., 2007). The integral of the strain energy
density within the fault system provides the internal work of the system:

$$W_{\rm int} = \iiint \frac{1}{2} \sigma_{ij} \varepsilon_{ij} dV$$
 Eq. 2

107 For a two-dimensional system in x-z space, Eq. 2 simplifies to:

108
$$W_{\rm int} = \frac{1}{2} \iint \sigma_{xx} \varepsilon_{xx} + \sigma_{zz} \varepsilon_{zz} + 2\sigma_{xz} \varepsilon_{xz} dx dz \qquad \text{Eq.3}$$

109 Although the internal strain energy represents elastic recoverable strain, the internal work 110 term also incorporates energy available for consumption by inelastic processes. For 111 example, pervasive deformation mechanisms, such as calcite twinning, may reduce stresses 112 within deformed host rock between faults and subsequently consume internal work. The 113 role of distributed permanent deformation on the work budget has not yet been 114 investigated. The strain energy density around faults is often greatest near fault tips and 115 irregularities. Consequently, the internal work is a powerful control on fault propagation 116 and many studies have predicted fault growth using minimization of internal work (Melosh 117 and Williams, 1989; Du and Aydin, 1992; Du and Aydin, 1996; Okubo and Schulz, 2005). *W*_{int} is reversible, meaning that it can either increase or decrease during fault system 118 119 evolution.

120 In the numerical implementation of *W*_{int}, the volume integral is calculated via 121 sampling points throughout the model. In the absence of faults, integrating a volume of 122 densely spaced sample points can provide accurate solutions for *W*_{int}; however, the 123 discretization of the fault into elements introduces approximations to the stress field. Near 124 fault tips, where stresses vary a lot over short distances, these approximations can impede 125 accurate sampling of *W*_{int}. Of all of the work components, calculation of *W*_{int} has the greatest 126 errors. Additionally, we need to consider that the total *W*_{int} is sensitive to the size of the

- 127 fault system analyzed. For two models with identical fault length, but different model
- 128 dimensions, the larger model will have greater *W*_{int}. For this reason, the difference in *W*_{int}
- 129 before and after fault growth can be more informative than its total value.

130 2.2 Work against gravity

131 Gravitational work considers the upward displacement, d_z , against gravitational 132 force, g,

$$W_{grav} = \iiint \rho g d_z(z) dV$$
 Eq. 4

134 where ρ is the density of the material. Like W_{int} , W_{grav} is reversible and can increase or 135 decrease during fault system evolution. Contractional fault systems will have an overall 136 upward displacement of points within the system that result in a positive W_{arav} while 137 extensional systems will have negative W_{grav} . Dempsey et al. (2012) show that a decrease 138 in W_{grav} can drive extensional faulting, even at the expense of an increase in W_{int} . This 139 complicates the elastic rebound paradigm of faulting, which was developed for strike slip 140 faults and assumes that slip events are associated with a decreased in stored elastic energy. 141 For all fault systems, decreases in the sum of W_{grav} and W_{int} , which is termed the 142 mechanical work, provide energy for fault slip, propagation and ground shaking. 143 The numerical calculation of W_{arav} requires sampling the displacement field 144 throughout the model. Unlike *W*_{int}, the displacements are not singular near the fault so 145 calculations of W_{arav} are not as sensitive to numerical discretization of the fault as 146 calculations of *W*_{int}. Both mechanical work terms are sensitive to the size of the model 147 domain.

148 2.3 Work against frictional resistance

The frictional work is the energy required to slide fault surfaces past each other. W_{fric} is irreversible, because the heat created during frictional slip exits the system. The onset of sliding involves processes related to the work required to generate new fault surfaces and W_{fric} (Fig. 2A). The component attributed to work against frictional resistance is calculated from the strength of the fault at stable sliding and the total slip, *s*. For a single 154 fault segment using a tension positive, compression negative sign convention, this155 component is calculated as:

$$W_{fric} = (c - \sigma_n \mu_d) sA = |\tau| sA$$
 Eq. 5

157 where *c* is cohesion, σ_n is normal stress, μ_d is the dynamic friction coefficient during sliding, 158 τ is the shear strength during sliding, and *A* is the slipped area of the fault. When positive, 159 σ_n is tensile along the fault and allows opening, so W_{fric} is zero. The complete frictional

160 work for one increment of loading is integrated over the surface area, *A*, of the fault so that:

161
$$W_{fric} = \oiint [c - \sigma_n \mu_d] dA = \oiint \tau dA$$
 Eq. 6

Frictional work can be calculated as fault strength evolves over a loading increment (e.g.
Savage and Cooke, 2010); however, most quasi-static studies use the stress state at the end
of the model convergence to calculate *W*_{fric} (e.g. Hardy et al., 1998; Burbridge and Braun,
2002; Cooke and Murphy, 2004; Del Castello and Cooke, 2008). Because frictional work is
inelastic, numerical implementations often load the faults system with multiple monotonic
steps so that *W*_{fric} is calculated for each loading step and integrated over the applied
loading (e.g. Cooke and Murphy, 2004; Del Castello and Cooke, 2008)(Fig. 2B).

169 Observations generally support the inference that mature faults have lower sliding 170 friction than immature faults (e.g. Marone, 1998). Nevertheless, the formulation of W_{fric} is 171 the same for the development of new fault surfaces and slip along mature fault surfaces. In 172 addition, a new fault surface with high strength and low slip may produce identical 173 frictional heating to a mature fault surface with low strength and high slip.

Until recently, we have not been able to measure the frictional heat produced by
fault slip events. On-going projects that drill and sample active fault zones (e.g. Townend et
al., 2009) and particularly efforts to drill soon after earthquake events (e.g. Fulton et al.,
2013) are providing constraints on frictional heating so that we can validate the theoretical
formulations. Studies of heating along exhumed faults are further constraining values of *W*_{fric} (Savage et al., 2014).

180 2.4 Seismic radiated energy

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181 The strength of a fault evolves during slip, from the initial strength (τ_0) to the 182 weakened sliding strength (τ) (Fig. 2A). This change in shear stress with slip produces 183 energy that is available for both fault propagation, W_{prop} , and ground shaking, W_{seis} , which 184 are irreversible and lost to the fault system. The portion of work allocated to each process 185 depends on the amount of fault slip relative to the slip required to bring the fault to its 186 sliding strength, referred to as the slip weakening distance, L (Fig. 2). All energy associated 187 with the shear stress drop ($\Delta \tau = \tau_0 - \tau$) when *s* exceeds *L* is available for ground shaking. In 188 other words, the energy required to generate the new fault surface is expended while $s < L_{i}$ 189 whereas work associated with s > L is available for ground shaking. If L = 0, W_{seis} for one 190 increment of loading is:

191
$$W_{seis} = \oiint \int \Delta \tau ds dA$$

192 Numerical models typically do not consider the evolution of shear stress with slip and only 193 calculate the change in shear stress, $\Delta \tau$, from comparing model results before and after 194 fault growth. In this case, the work need not be integrated over slip and Eq. 7 simplifies to:

195
$$W_{seis} = \oiint \frac{1}{2} \Delta \tau s dA$$
 Eq.8

196 In the event of multiple loading increments, W_{seis} is calculated from the change in shear 197 stress and the change in slip from one loading step to the next (Fig. 2B). The calculation of 198 the W_{seis} is not largely affected by discretization.

199 The formulation of W_{seis} is the same for both established and new fault surfaces, but 200 the values of τ and τ_0 may differ. For example, intact rock often requires greater shear 201 stress to reach failure than an established fault surface. While quasi-static models cannot 202 be used to calculate shaking directly, they can be used to examine $\Delta \tau$ and estimate the 203 energy available for shaking.

204 2.5 Energy of fault propagation/growth

205The energy of fault propagation, W_{prop} , is the energy required to create206discontinuities within a material. This energy is irreversible within most natural207conditions. Although Obriemoff (1930) was able to reverse the growth of a crack in mica by208growing the crack within a vacuum, under atmospheric conditions, available cations209bonded to new crack surfaces and impeded healing.

Eq. 7

Griffith (1921) demonstrated that the energy release rate, *G*, required to grow a crack can be calculated directly from the stress intensity factors on the crack, *K_i*, where *i* refers to growth by modes I, II or III:

213
$$G = \frac{1 - v^2}{E} \left[K_I^2 + K_{II}^2 + \frac{K_{III}^2}{(1 - v)} \right]$$
 Eq. 9

Here, *v* is Poisson's ratio and *E* is Young's Modulus. However, this formulation presumes
co-linear crack propagation of a single discontinuity, so Eq. (8) works well for openingmode propagation of joints and veins in rock, and is applicable for controlled experiments
that result in co-linear growth. However, this formulation does not apply universally to
faults, which often propagate out of plane. Consequently, predicting fault growth using *G* is
problematic.

220 Taking an observational approach, some workers have estimated W_{prop} from 221 measurements of damage around faults. Using the measured energy release rate of 222 opening mode cracks ($G \cong 1 \text{ J/m}^2$), the total energy consumed in fault growth can be 223 estimated by summing the damage area around faults (e.g. Wong, 1982, 1986; Cox & 224 Scholz, 1988). For cases of new faulting within intact rock, these values are 10^{5} - 10^{6} J/m² 225 (Wilson et al., 2005; Pittarello et al., 2008). In contrast, values from laboratory samples 226 range only up to 10⁴ J/m² (Wong, 1982, 1986; Cox & Scholz, 1988; Lockner et al., 1992). 227 These different ranges from laboratory and field measurements highlight the need to 228 develop a better understanding of W_{prop} . In section 4, we present numerical models that 229 simulate laboratory experiments of faults loaded at values critical for growth. In addition 230 to demonstrating the work budget of fault growth, these experiments allow for calculation 231 of W_{prop} . For co-linear propagation, this value can be validated against calculations of G.

A formula for W_{prop} that does not require co-linear growth is related to that for W_{seis} . As mentioned in the previous section, the portion of work allocated to each of these processes depends on the amount of fault slip, *s*, relative to the slip weakening distance, *L* (Fig. 2A). The energy required to generate the new fault surface is expended while *s* < *L*, whereas work associated with *s* > *L* is available for ground shaking. For loading steps where the total slip remains less than L, this is:

238
$$W_{prop} = \oiint \frac{1}{2} \Delta \tau \Delta s dA \qquad \text{Eq. 10}$$

where Δs is the change in slip associated with that loading step and W_{seis} equals zero. For loading steps that produce slip that spans the slip weakening distance, such as step 3 on Fig. 2B, this is:

242

$$W_{prop} = \oiint \frac{1}{2} \Delta \tau s_L dA$$
 Eq. 11a

243

$$W_{seis} = \oiint \frac{1}{2} \Delta \tau (\Delta s - s_L) dA$$
 Eq. 11b

Here, s_L represents slip during this loading step when s < L. The total W_{prop} and W_{seis} for the system are summed at each loading step.

246 2.6 External work

The external work reflects the overall mechanical efficiency of the fault system, such that an efficient system will require less W_{ext} to accommodate the same tectonic displacement than an inefficient system. W_{ext} is the product of the shear and normal tractions, τ and σ_n , and shear and normal displacements, u_s and u_n along the model boundaries, B (Fig. 1):

252
$$W_{ext} = \oiint (\tau u_s + \sigma_n u_n) dB$$
Eq. 12

If multiple loading steps are applied, the calculation of W_{ext} should be integrated over the loading steps. The value of W_{ext} also equals the sum of the energy budget components in Eq. (1). This relationship provides an independent check on the calculations of individual energy terms within the budget. Considerations of fault system evolution benefit from attention to changes in external work, ΔW_{ext} , between stages of fault growth. To meet requirements of energy conservation, we can consider that ΔW_{ext} equals the sum of the associated changes in all work budget components.

This energy budget differs from the energy considerations single earthquakes, where it is assumed that no change in remote loading or strain occurs during the short time-span of fault slip (e.g. Scholz, 2002). Under this assumption, which follows that presented by Griffith (1921), the change in mechanical work, $\Delta(W_{int} + W_{grav})$, equals the sum of the changes in W_{fric} , W_{prop} and W_{seis} . During an earthquake, the decrease in mechanical work provides the energy that is consumed in frictional heating, generation of damage, and ground shaking. The assumption that $\Delta W_{ext} = 0$ is not appropriate for studies 267 of fault evolution, however, which consider fault development over multiple earthquake 268 cycles due to tectonic loading. Consequently, we consider ΔW_{ext} as the fault system evolves, 269 following the work of Del Castello and Cooke (2007), which shows that a fault system 270 prefers to propagate in the direction that maximizes ΔW_{ext} .

271 The approach of analyzing the increasing efficiency of evolving fault systems in 272 numerical models depends on the method of applied deformation, i.e. applied tractions or 273 displacements. For fault systems that are loaded with displacement along the model 274 boundaries, fault growth will increase the compliance of the system, so that the 275 corresponding tractions along the model boundaries decrease. In this case, *W*_{ext} will 276 decrease with the growth of more efficient faults within the system. In contrast, if the 277 system is loaded with tractions along the model boundaries, then the increase in 278 compliance due to fault growth will increase the displacement of the model boundaries. In 279 this case, *W_{ext}* will increase with the growth of more efficient faults within the system. 280 Either displacement or traction loading is suitable for a work budget analysis, but we 281 advise against using mixed non-zero boundary conditions. In the cases of both 282 displacement only and traction only loading, the most efficient fault growth path is that which maximizes the change in external work, ΔW_{ext} , from before to after fault growth. We 283 284 demonstrate both types of loading in section 4, but continue to refer to growth that 285 optimizes energy efficiency as "work minimization".

286 **3. Work minimization in structural geology**

287 For over a century, engineers have used energetic principles to understand material 288 deformation and failure. Griffith (1921, 1924) used the energy budget associated with 289 opening mode fracture growth to determine the equations governing fracture propagation, 290 which launched the field of fracture mechanics. In an interesting revisitation of this classic 291 approach, Engelder and Fisher (1996) execute Griffith's conceptual model within the 292 laboratory and discuss implications for the loading of joints in the crust. Here, we review 293 past applications of energy minimization to assess fault growth in detail. These studies 294 present many interesting insights, as well as challenges that remain to be resolved. 295 Though we do not review them in detail, work minimization approaches have been

296 used in other geoscience disciplines as well. For example, within metamorphic petrology, 297 the numerical code ELLE predicts microstructure development based on work 298 minimization of forces and chemical reactions within the system (e.g. Jessell et al, 2001). 299 Modern fluvial geomorphology is founded on the work of Langbein and Leopold (1964), 300 which analyzes the energy budget of evolving river channels. Another formulation of work 301 minimization is maximum entropy. Rather than assessing the energy available within the 302 system, the maximum entropy approach assesses the irreversible energy that is no longer 303 available to the system. An example of this application is atmospheric circulation models 304 that utilize maximum entropy to predict flow (e.g. Kleidon et al., 2003).

305

306 *3.1 Minimum viscous dissipation*

307 For several decades, the minimization of work has been utilized to predict the 308 development of fault systems. The earliest approaches minimized the viscous dissipation 309 within ocean ridge and transform fault systems. In this application, analytical solutions 310 characterizing the system were solved to find the fault network that uses the least force to 311 accommodate the prescribed plate boundary displacements via viscous flow within the 312 fault zones. Lachenbruch and Thompson (1972) showed that minimum viscous dissipation 313 predicts the orthogonal orientation of ridges and transform faults. This study was followed 314 by several others that augmented this analysis to successfully account for various 315 observations of ocean ridge geometry (Fig. 3) (e.g. Froidevaux, 1973; Stein, 1978; Kleinrock 316 and Morgan, 1988). In 1979, Bird and Yuen published a critique of the approximations used 317 in the analytical implementation of the minimum viscous dissipation approach. They 318 pointed out that the approximations are avoided with numerical implementations that 319 directly solve the governing equations and consider more complex rheology and coupled 320 processes within the crust (e.g. Regenauer-Lieb et al., 2006; Holtzman et al., 2005). This 321 critique was followed by a comment in support of the original methodology by Sleep et al. 322 (1979).

A similar approach to minimizing viscous shear is minimizing on-fault shear stress, which has been used to predict the orientation of frictionally sliding faults. Reches (1983) 325 implemented a minimum dissipation approach for frictionally sliding faults, which showed

326 that sets of orthorhombic faults require the least shear stresses to accommodate three-

327 dimensional states of strain. In this same study, Reches (1983) demonstrated that the fault

328 networks that minimize shear stress also minimize the internal work in the rock between

329 the sliding surfaces.

330 *3.2. Minimization of internal work*

331 More efficient fault systems allow for more fault slip, so that the surrounding rock is 332 relatively less strained and subject to less stress. Consequently, the minimization of 333 internal work, which integrates the product of the stress and strain fields Eq. (2), can 334 achieve similar predictions of fault geometry as methods that minimize the forces required 335 to accommodate prescribed deformation. Melosh and Williams (1989) demonstrated that 336 antithetic pairs of normal faults result in lower internal work than synthetic pairs of 337 normal faults. The greater efficiency of antithetic normal faults may account for the 338 frequent occurrence of antithetic normal faults within grabens (Melosh and Williams, 339 1989). Hexagonal cracks within uniformly expanding materials, such as mud cracks, are 340 another common observation that has been explained by the minimization of strain energy 341 per unit of new crack area (Lachenbruch, 1962).

342 Within elastic systems, the spatial variations of internal work throughout a fault 343 system can be mapped with the Strain Energy Density, SED, so that SED is the internal 344 work at a point within the system. The distribution of SED has been used to assess the 345 growth and propagation faults. This approach utilizes the premise that faults develop in 346 regions of high stress and strain, so that the growth of faults should minimize the total SED, 347 i.e. the internal work (e.g. Du and Aydin, 1993; Du and Aydin, 1996; Okubo and Schulz, 348 2005; Olson and Cooke, 2005). Following this approach, Olson and Cooke (2005) used 349 three-dimensional models to compare fault orientations predicted by the locations of high 350 SED with those predicted by Coulomb failure theory, and showed that locales of high SED 351 better predict the interpreted evolution of imbricate thrust faults within the Los Angeles 352 basin (Fig. 4). Griffith and Cooke (2004) used SED patterns around alternative three-353 dimensional configurations for the intersection of the Whittier and Puente Hills faults to

show that the configuration that minimizes internal work also reproduces the observedslip rate.

356 Further insights can be gained by decomposing SED into its two deformational 357 components, the dilatational and distortional strain energy densities (e.g. Jaeger et al., 358 2007). Because faults accommodate distortion, the pattern of faulting may be predicted by 359 the maximization of distortional SED (e.g. Du and Aydin, 1992; Du and Aydin, 1996; Okubo 360 and Schulz, 2005). Limiting analysis to the distortional SED is akin to using a maximum 361 shear stress or second invariant of shear stress failure criterion, such as von Mise's 362 criterion. While this approach does not consider the role of the normal stress, as Coulomb 363 failure does, the predicted fault propagation patterns resemble general observations of 364 fault step-overs (Du and Aydin, 1993) and the propagation path of deformation bands (Okubo and Schulz, 2005), as well as the development of the Eastern California Shear Zone 365 366 north of a bend in the San Andreas fault in California, USA (Du and Aydin, 1996),. 367 Furthermore, Okubo and Schulz (2005) used dilatational SED to predict fault initiation 368 locations, which may be related to dilatational failure, in contrast to distortion dependent 369 propagation processes. These studies demonstrate that faults grow into regions of high 370 internal work, which consequently can reduce the W_{int} of the system.

Another application is in crustal models, where minimum internal work is used to interpolate deformation in regions where geologic data are scarce (e.g. Saucier and Humphries, 1993; Peltzer and Saucier; 1996). These models use geodetic velocities and geologic slip rates as input where they are available. Minimizing internal work provides a way to resolve slip along portions of the faults, while satisfying data constraints. The widespread use of this approach supports the premise that faults develop to minimize internal work.

A corollary to the minimization of internal work is the maximization of slip. The theory here is that the fault system with the least internal work will display the largest amount of fault slip. DeBremaecker and Ferris (2004) show that maximum slip predicts the orientation of wing-crack development at fault tips. Using analog experiments of restraining bends, Cooke et al. (2013) showed that the percentage of slip accommodated by fault systems increases with the propagation of new faults as the system becomes more

efficient. This analog modeling study further supports the premise that faults evolve togreater efficiency by maximizing fault slip and minimizing internal work.

386 3.3 Minimization of strain

387 The calculation of either internal work or viscous dissipation requires solving for 388 both the displacement (or velocity) and the stress (or stress rate) fields. An analytically 389 more simple and computationally faster approach is to minimize only the strains or 390 displacements. By assuming that locations of high stress are also locations of high strain, 391 minimizing strain should have the same effect as minimizing internal work. The 392 computational benefit is that the kinematic solutions can be used to find the strain field, 393 without having to solve for stress. This approach has been used to assess the ratios of pure 394 shear to simple shear during non-steady state deformation (Fossen and Tikoff, 1997). 395 Following publication of that paper, Jiang (1998) criticized the minimum strain path 396 approach, because it did not consider the full energetics of the system. In their reply, 397 Fossen and Tikoff (1998) note that, due to the complexities of the system, it is not known if 398 minimizing work gives a different result from the kinematic solutions found by minimizing 399 strain. In many ways, this discussion and reply echo the discussion two decades prior 400 between Bird and Yuen (1979) and Sleep et al. (1979).

Minimization of strain also is used in several applications to determine fault slip, when the system is under-constrained. Strain minimization can be used in a similar way to the minimization of internal work, to resolve unconstrained fault slip rates within crustal deformation models (e.g. Flech et al., 2001). Similarly, several three-dimensional structural restoration algorithms minimize strain to determine slip along faults within a system (e.g. Plesch et al., 2007; Shackleton et al., 2009).

407 *3.4 Limit analysis*

Limit analysis provides an alternative approach for finding the most efficient failure
surface using analytical solutions. Like other methods, it draws from engineering
mechanics and has a long history of success. For example, Bishop (1955) developed a
standard method to investigate slope stability by delineating the most likely landslide

412 surface as that which fails at the least applied stress. Similar methodology has been used to 413 find the fault surface within an accretionary wedge that fails under the lowest force (e.g. 414 Maillot and Leroy, 2003; Maillot and Leroy, 2006; Cubas et al., 2008; Souloumiac et al., 415 2009; Souloumiac et al., 2010; Mary et al., 2013). The limit analysis requires analysis of 416 many potential surfaces to find the one that requires the minimum force to deform. This 417 surface is at its failure limit, so it requires the 'least upper bound of applied force', a phrase 418 often used in this literature. Lesser force can be applied, but the upper bound of force is 419 that required for failure of the system, i.e. the development of faults. One of the great 420 benefits of the limit analysis approach is that it is analytical, so that the search for the 421 preferred surface completes within seconds. Consequently, limit analysis can be used to 422 assess sequential failure of multiple surfaces and shows many of the same features as 423 accretionary wedges (Fig. 5) (e.g. Cubas et al., 2008, Mary et al., 2013).

424 3.5. Minimization of total and external work

425 Many workers have recognized that internal work does not capture the complete 426 system and so have sought to find the fault geometry that minimizes the system's total 427 work. This total work represents different combinations of work budget components in 428 different applications, including: 1) work against gravity, which is sometimes integrated 429 with internal work, 2) work against friction, 3) energy required to propagate faults, and 4) 430 seismic radiated energy (Fig. 1) (e.g. Cooke and Murphy, 2004). Jones and Wesnousky 431 (1992) considered internal work and work against friction to demonstrate that under 432 transpressional loading, slip partitioning between two parallel fault planes, one with dip 433 slip and another with strike slip, minimizes the total work of the deforming system. Cooke 434 and Kameda (2002) showed that, for a variety of proposed cross-sections of the Los 435 Angeles basin, the system that requires the least total work to accommodate tectonic 436 loading is that which best matches the available fault slip rates. In a study of the Lake 437 Meade Fault zone, Marshall et al. (2010) assessed the relative efficiency of different fault 438 interpretations. The configuration that minimizes work is consistent with independent 439 observations and provides inside into interaction among faults. These studies lend support 440 to the postulation that fault systems evolve to minimize total work.

441 By far, the bulk of investigations of fault evolution using work minimization with a 442 full work budget analysis examine accretionary wedge systems. Both analog models and 443 crustal wedges provide excellent laboratories for studying fault growth, because the 444 material is relatively undeformed prior to incorporation into the wedge, while within many 445 other deformational regimes, the rock mass may have inherited fault structures that 446 complicate a minimum work analysis. Dahlen et al. (1984) is a watershed study that 447 introduced the concept of critically tapered wedges. A few years later, Mitra and Boyer 448 (1986) framed the development of duplexes, which can be viewed as the cellular unit of an 449 accretionary wedge, in terms of the system's work budget. Dahlen et al. (1989), followed up 450 with a work budget for the entire accretionary wedge that has been augmented by many 451 workers (e.g. Gutcher et al., 1998; Hardy et al., 1998; Burbridge and Braun, 2002).

452 Several studies assess fault propagation within accretionary systems through 453 minimization of the energy required to overcome friction and gravity (Gutscher et al., 1998; 454 Hardy et al., 1998; Burbridge and Braun, 2002). Gutscher et al. (1998) showed analytically 455 that minimization of frictional and gravitational work accurately predicts the length of a 456 new accretionary forethrust. Burbidge and Braun (2002) extended this analysis to 457 determine the friction required for frontal and back accretion, while Hardy et al. (1998) 458 used a combined Eulerian-Lagrangian scheme within numerical models to minimize 459 frictional and gravitational work to predict the progressive sequence of accretion. Masek 460 and Duncan (1998) expanded these analyses by including internal work, showing that this 461 component has a large influence on fault geometry. As with previous applications of work 462 minimization, Masek and Duncan (1998) evoked a discussion and reply (DeBremaecker, 463 1999; Masek and Duncan, 1999). However, this discussion focuses on the effect of 464 numerical discretization on the work budget, rather than on the premise of a fault growing 465 to minimize work.

In a recent study of accretion, Del Castello and Cooke (2007) used a numerical
implementation of work minimization to simulate analog experiments of accretion and
present the evolving work budget during the process of new forethrust development (Fig.
6). In these finite models, the external work acting on the system represents the system's
total work, and equals the sum of the five work budget components. A new forethrust
reduces the external work required to deform the system due to a reduction in frictional

472 work, even though the new fault increases the system's internal work (Fig. 6B).

- 473 Additionally, the investigation of potential position and vergence of thrusts shows that the
- 474 thrust in the numerical model that minimizes W_{ext} matches the forethrust that develops in
- 475 the experimental sandbox. Del Castello and Cooke (2007) further found that propagation
- 476 of the forethrust occurs when the decrease in external work due to adding the fault (ΔW_{ext})
- 477 exceeds the work required to create the new fault surface ($W_{prop} + W_{seis}$; Fig. 6C). This
- 478 outcome highlights the critical role of the work of fault propagation, a work term not
- 479 previously considered. This study suggests that consideration of external work and the
- 480 work of fault propagation together allow for prediction of not only the path of fault
- 481 propagation, but also the timing of fault propagation.

482 **4. Work minimization predictions of fault propagation**

483 To demonstrate the power of using a work minimization approach to study fault 484 growth, we present the propagation paths and work budgets for co-linear growth of a joint 485 and faults both with and without slip weakening. Following investigations by Del Castello 486 and Cooke (2007), the most efficient growth paths are found by maximizing ΔW_{ext} , the 487 difference in *W_{ext}* before and after propagation. Potential growth elements, or *pupative* 488 *elements*, are radial to the fault tip at a prescribed range of angles to the fault. The growth 489 of the fault via each pupative element is assessed separately and comparison of the results 490 reveals the most efficient radial propagation path. The stresses along some pupative 491 elements will not meet the tensile or shear failure criterion, so the fault will not be able to 492 grow in these directions and W_{ext} remains unaltered (i.e. $\Delta W_{ext} = 0$). Among the pupative 493 elements that do fail, the element that maximizes ΔW_{ext} is added to the initial structure. 494 Though only one growth iteration is presented here, this sequence can be repeated for 495 subsequent pupative elements.

We simulate joint and fault propagation under the conditions of laboratory
experiments on fault growth by Bobet and Einstein (1998). Both the joint and the fault are
12.7 mm long and are comprised of 100 boundary elements of 0.127 mm embedded within
a slab of gypsum (Fig. 7A, 7B). The pupative elements have the same length as elements
along the pre-existing fault. The orientations of pupative elements are considered at 5°

angle increments between 45° and 315° clockwise from the joint or fault tip (Fig. 7B inset).
We use the material properties for gypsum from Bobet and Einstein (1998) with a water to
gypsum ratio of 0.4, tensile strength of 3.2 MPa, an average Poisson's ratio of 0.15, and an
average Young's modulus of 5.96 GPa.

505 We utilize the two-dimensional boundary element method (BEM) program Fric2D 506 (Cooke and Pollard, 1997). Dislocation surfaces made up of a series of elements of equal 507 length are free to open or slip, but not to interpenetrate, in response to the tractions or 508 displacement applied on model boundaries and interactions with other elements. Fric2D 509 solves the quasi-static equations of deformation on all elements. We take a compression 510 negative, tension positive sign convention. Elements fail in tension when the normal stress, 511 σ_{p} , exceeds the prescribed tensile strength, T:

512 $T \ge \sigma_n$.Eq. 13513Shear failure is governed by a frictional failure criterion, so that an element along an514established fault slips when the shear stress, τ , exceeds the frictional strength, which is the515sum of cohesion, c, and the product of the coefficient of static friction, μ_s , and σ_n :516 $|\tau| \ge c - \mu_s \sigma_n$ 517Fric2D also captures slip-weakening behavior (Savage and Cooke, 2010). When slip on a

518fault element exceeds a prescribed weakening distance, *L*, the friction evolves linearly from519its static value, μ_s , to a dynamic value, μ_d .

520 The criterion governing the shear failure of intact rock has the same form as Eq. 521 (14), but *c* is replaced by the rock's inherent shear strength, S_o , and μ_s is replaced by the 522 coefficient of internal friction, $\mu_{o:}$

523 $|\tau_{o}| \geq S_{o} - \mu_{o} \sigma_{n}$

Eq. 15

Following onset of failure in intact rock, laboratory tests have shown that fault strength decreases as a fault surface develops during slip (e.g. Handin and Hager, 1957; Hazzard and Young, 2000; Mitchell and Lockner, 2008). In Fric2D, when a pupative element fails, its strength decreases so that shear stress drops from τ_0 to τ as its internal friction value, μ_0 , drops to μ_d . The inelastic behavior of frictional slip is solved iteratively until the model converges (Cooke and Pollard, 1997).

530 4.1 Joint propagation

In the joint model, we apply horizontal and vertical displacement boundary conditions, respectively u_h and u_v , on the edges of the gypsum block (Fig. 7A). Keeping $u_h=0$, we adjust u_v to its critical value, when the stress intensity factor, K_I , at the joint tip equals gypsum's mode I fracture toughness, K_{Ic} , of 0.15 MPa \sqrt{m} . This value of K_{Ic} value is from the curve reported by Chen et al. (2006, Fig.2), extrapolated to the 0.4 water to gypsum ratio used by Bobet and Einstein (1998). The failure criterion is met at $u_v=2.71e^{-5}$ m, which produces ~ 3.1 MPa of axial tension on the model boundaries.

538 Assessment of pupative elements around the joint tip reveals that the most efficient 539 orientation for growth, the orientation that minimizes ΔW_{ext} , is at 180° (Fig. 8a). This co-540 linear growth is commonly observed for joints under perpendicular extension (e.g. Pollard 541 and Aydin, 1988). *W_{ext}* is 1.1135 J before growth and drops to 1.1131 J after growth, 542 resulting in ΔW_{ext} per pupative element area of -3.63 J/m². Note that, though the model is 543 two-dimensional plane strain, all elements have a width of w = 1 m out of the plane under 544 consideration. The components of the joint's work budget are summarized in Table 1 and 545 Fig. 9. For this growth scenario, W_{seis} , W_{fric} and W_{grav} are 0, and $\Delta W_{int} = \Delta W_{ext} = W_{prop}$. Using 546 Eq. (10), we find that $W_{prop} = 3.62 \text{ J/m}^2$, which balances with ΔW_{ext} , indicating that the 547 change in external and internal work is a result of energy spent on propagating the joint. 548 Because the joint growth is co-linear, we can compare this estimate of W_{prop} to that estimated by the critical energy release rate, G_c . Using $K_{lc} = 0.15$ MPa \sqrt{m} , we find from Eq. 549 (9) that $G_c = 3.71 \text{ J/m}^2$. Thus, $\Delta W_{ext} \cong W_{prop} \cong G_c$, demonstrating consistency between 550 551 opening-mode joint growth by work minimization and the theory outlined by Griffith 552 (1921) for co-linearly propagating cracks.

553 4.2 Fault Propagation (Fault 1)

554 For the more complex scenario of fault growth, we began with two widely spaced, 555 60° dipping faults in gypsum that demonstrate co-linear growth under biaxial loading 556 (scenario 60°-2a-4a from Bobet and Einstein (1998)). While Bobet and Einstein (1998) 557 suggest that faults with internal tips at least 3 half-lengths (19.05 mm) from one another 558 behave as isolated faults, the numerical simulations show some interaction between the stress fields surrounding these faults, even at 28.40 mm of internal tip spacing. To remove
the effects of any interaction, we include only one fault in these simulations of fault growth
(Fig. 7B). While this geometry results in a slight overestimate of work budget components,

- 562 it provides a more robust demonstration of isolated fault growth patterns under biaxial
- 563 loading.

564 The numerical model for Fault 1 contains a pre-existing fault of 12.7 mm dipping 565 60°, with zero cohesion and equal static and sliding friction coefficients ($\mu_s = \mu_d = 0.3$) (Bobet, 2000). The coefficient of internal friction on potential growth elements, μ_o , also is 566 567 0.3. The inherent shear strength, $S_0 = 78.7$ MPa, for the growth elements is found by solving 568 Eq. (15) using the critical shear and normal stresses on the element oriented at 180° from 569 the fault tip at the onset of fault propagation (Fig. 7B) (Bobet, 1997, Table 4.2; Bobet and 570 Einstein, 1998). We apply traction boundary conditions on the horizontal and vertical 571 edges of the gypsum block of σ_h = -2.5 MPa and σ_v = -29 MPa, respectively (Bobet and 572 Einstein, 1998).

573 ΔW_{ext} is largest for the pupative growth element oriented at 180° from the fault tip 574 (Fig. 8b), consistent with the experimental observations of co-linear growth. The complete 575 work budget for Fault 1 is compared with that for the joint in Fig. 9 and Table 1. For growth 576 of 1 element length of 0.127 mm, W_{ext} increases from 878.5522 J to 878.5947 J. In contrast 577 to the joint model, which used displacement boundary conditions so that a decrease in W_{ext} 578 indicates an increase in efficiency, in this model an increase in W_{ext} indicates an increase in 579 efficiency (see discussion in Section 2.6). This results in a ΔW_{ext} per pupative element area 580 of 333.83 J/m². For this scenario, W_{seis} and W_{grav} are zero and $\Delta W_{ext} = \Delta W_{int} + \Delta W_{fric} = W_{prop}$ 581 + ΔW_{fric} . Using Eq. (6) and Eq. (10), we find that $W_{fric} = 79.78 \text{ J/m}^2$ and $W_{prop} = 260.77 \text{ J/m}^2$. 582 The sum of W_{fric} and W_{prop} is 340.54 J/m², which balances with ΔW_{ext} , indicating that the 583 change in external work results from energy spent on fault propagation and the increase in 584 frictional sliding that fault growth allows.

585 As for the joint, this co-linear propagation allows for a comparison of the fault's 586 W_{prop} with that estimated by the critical energy release rate, G_c . Using the fracture 587 toughness reported by Fric2D at the fault tip at critical loading ($K_{IIc} = -1.27 \text{ MPa}\sqrt{\text{m}}$), we find that the critical energy release rate is $G_c = 263.1 \text{ J/m}^2$. Thus, G_c is consistent with the work required for fault growth calculated using a work budget approach.

590 This model demonstrates the robustness of using a work minimization approach to 591 model fault growth. Because this approach considers the global ΔW_{ext} instead of the local 592 *K*_{IIc}, it can be applied to more complex fault propagation situations and is not restricted to 593 co-linear growth. The work minimization approach also provides an estimate of the energy 594 required for breaking intact rock to form new fault surfaces. In this case, we find that W_{prop} 595 for fault growth in gypsum is 260.77 J/m^2 . This value is reasonable, at two orders of 596 magnitude larger than W_{prop} for joint growth in gypsum and several orders of magnitude smaller than the work inferred to grow faults in rock $(10^4-10^6 \text{ J/m}^2)$ (Wong, 1982, 1986; 597 Cox & Scholz, 1988; Lockner et al., 1992; Wilson et al., 2005; Pittarello et al., 2008). 598

599 4.3 Propagation of slip-weakening faults (Fault 2, Fault 3)

600 We consider slip-weakening along pupative growth elements in two additional 601 scenarios, Fault 2 and Fault 3. We prescribe $\mu_o = 0.6$ and $\mu_d = 0.3$ along both faults, but vary 602 the weakening distance, *L*, over which the drop in friction from μ_o to μ_d occurs. L = 4x10⁻⁶ m 603 for Fault 2 and L = 1x10⁻⁶ m for Fault 3. The shear stress drop resulting from slip 604 weakening has the potential to produce seismic radiated energy, *W*_{seis} in addition to *W*_{prop} 605 (Fig. 2).

As shown in Table 1 and Fig. 9, varying *L* does not change the angle of the pupative element that maximizes efficiency, which remains at 180° for both Fault 2 and Fault 3, nor does it change ΔW_{int} or ΔW_{fric} . However, *L* does affect the partitioning of work between fault propagation, W_{prop} , and the production of seismic waves, W_{seis} (Fig. 2). Shorter *L* (Fault 3) means that less energy is required to break the intact rock ahead of the fault, leaving more energy for ground shaking.

612 4.4 Comparison of work minimization and the Coulomb criterion

613 Correspondence between the maximum mechanical efficiency prediction of a fault's
614 propagation path with that observed in laboratory experiments supports the use of work
615 minimization as a fault growth criterion. This implementation of the principle of work

616 minimization using numerical models utilizes Coulomb theory as a failure criterion for both 617 the potential growth elements extending from the fault tip and those along the pre-existing 618 fault surface. However, this approach differs from Coulomb stress analyses that determine 619 a fault's propagation path from planes of maximum Coulomb shear stress out ahead of a 620 fault tip. We compare the two approaches here. 621 The Coulomb shear stress is found on a potential failure surface by rearranging Eq. 622 (15) to isolate S_{o} , replacing it with the Coulomb stress, σ_{c} , and replacing the frictional shear

623 strength, $τ_0$, with the shear stress resolved on the surface, τ:

624

625 Positive σ_c along a surface indicates that it has the potential for failure, while negative σ_c 626 indicates that failure is inhibited. σ_c is useful for finding the relative failure potentials of 627 different fault surfaces.

628 The potential failure planes that maximize σ_c are at angles of $\pm \varphi$ to the maximum 629 compressive stress, σ_3 . φ depends upon the inherent friction coefficient, μ_o :

630

 $\pm \varphi = \frac{1}{2} \operatorname{atan}(1/\mu_0)$ Eq. 17

631 The \pm indicates that Coulomb theory predicts two potential planes, on either side of σ_3 , that 632 carry the same shear stress magnitude and therefore the same failure potential. Selecting 633 between these planes is not trivial.

634 For this analysis, we consider both Coulomb and tensile failure near the fault tip. Fig. 635 10A shows the orientations of planes carrying the maximum tension and Coulomb stress, . 636 Figs. 10B and 10C display the contours of the maximum tensile stress and Coulomb stress, 637 respectively, following fault slip and prior to fault propagation. In Fig. 10B, tensile stress is 638 largest below and to the right of the fault tip, within the tensile quadrant of this left-lateral 639 fault. Tension values exceed the tensile strength of 3.2 MPa in much of this region. Fig. 10C 640 shows that the maximum Coulomb stress is located out ahead of the fault tip. The Coulomb 641 stress fails to exceed the inherent strength of the gypsum ($S_0 = 78.7$ MPa) in much of this 642 region, except for directly in front of the fault. If we focus in the vicinity of growth, at half of 643 1 element length (0.0635 mm) from the fault tip (Fig. 10D), the Coulomb stress reaches 644 87.1 MPa and exceeds the shear strength of the rock. In addition, $\sigma_1 = 72.3$ MPa at this 645 location, exceeding the tensile strength of the gypsum (T = 3.2 MPa). T is exceeded at the 646 two locations to the right of the fault tip as well, where $\sigma_1 = 82.2$ MPa and 75.4 MPa.

647 Predicting the propagation path from these results is problematic. First, tensile 648 stresses are large and exceed the tensile strength of the rock in multiple locations. This 649 suggests that tensile failure in the form of a wing crack off the tip of the fault would occur 650 prior to any failure in shear. Second, ignoring the potential for tensile failure, both Coulomb 651 planes at the location where the shear strength is exceeded out ahead of the fault tip have 652 equal potential for failure. The plane with left-lateral slip aligns with the pre-existing fault 653 and may be the better candidate for fault growth in this scenario, but such a determination 654 is difficult to automate, as it may not be possible in other scenarios. Furthermore, selecting 655 incorrectly will greatly alter the prediction of any subsequent growth.

656 Using a work minimization approach has the advantage of evaluating both tensile and shear failure by another metric, in this case the energy efficiency of the system, and 657 658 consequently avoids this fault plane ambiguity. The tensile failure of wing cracks is 659 considered explicitly within the work minimization analysis, but under the loading used 660 here, wing cracks are not as efficient as in-plane fault propagation. Considering the tensile 661 failure and Coulomb failure independently at points ahead of the fault does not provide 662 insights in the competition between these two mechanisms. In addition, within the work 663 minimization approach, a pupative growth element is evaluated as a continuation of the 664 pre-existing fault and experiences the slip and stress associated with that configuration. 665 For many scenarios, this provides an advantage over evaluating slip at a point ahead of the 666 fault tip. While the process zone around the fault tip no doubt includes microcracks that 667 are not radial to the fault tip, consideration of tip-radial propagation conveniently 668 considers the overall fault propagation path.

669 **5. Implications**

Examination of the complete energy budget provides insight into the competing
influences of different processes within deforming fault systems. In an innovative study of
multilayer folding, Ismat (2008) uses work minimization to explore the competition
between different flexural mechanisms within the competent and incompetent layers of an
evolving fold. Because gravity controls the growth of accretionary wedges, we might
expect gravitational work to be the largest part of the accretionary experiment energy

676 budget. DelCastello and Cooke (2007) showed that, within sandbox models of accretion, the 677 greatest energy is consumed by the work of against friction. The frictional work arises from 678 the weight of the overlying material, so gravity also plays a role in this scenario. The 679 growth of new faults in front of the wedge serves to decrease frictional work, at the 680 expense of an increase in internal work. The full work-budget analysis undertaken by 681 DelCastello and Cooke (2007) revealed this complex trade-off between different 682 components of the work budget. At a much larger scale, Meade (2013) showed that, within 683 crustal scale models, internal work is the dominant energy sink. Thus, different systems 684 require attention to different components of the work budget.

685 One challenge for any work budget investigation of fault systems is the appropriate 686 bounds for the system. Consider a complex network of active faults, such as at the 687 boundary of the Pacific and North American plates in southern California. Perhaps this 688 system has evolved to minimize work at the plate boundary scale, but can we extract 689 portions of this system (e.g. the Los Angeles basin) and reliably assume that this portion 690 should also minimize work? It could be that local mechanical inefficiencies develop within 691 a system that is efficient overall. If the system under consideration within the model is too 692 constrained in scale, then we may not be allowing for local inefficiencies to arise. Another 693 way to think of this is that we do not yet know if work minimization predictions of fault 694 evolution are independent of scale.

695 Another challenge in work minimization analyses is that active fault systems do not 696 always realize the most efficient configuration. We see this within sandbox models of 697 accretion, where underthrusting persists until a new fault can grow, even though a new 698 fault would be more efficient much earlier in the deformation (Fig. 6) (DelCastello and 699 Cooke, 2007). However, growth of this new thrust fault is not possible until the system has 700 enough energy to form the new fault surface. This energy requirement suggests that the 701 degree of inefficiency that a fault system can tolerate depends on the strength of the crust, 702 as well as on the details of the fault network geometry. It may be that inefficient fault 703 configurations persist longer in regions of thick-skinned tectonics, where crustal strength 704 is high, than in accretionary systems, where material is weaker.

We now have the tools and methods to make great advances in the study of faultsystem evolution via work minimization. In the application demonstrated here using

707 numerical simulations of laboratory experiments, minimum external work determines the 708 direction of the most efficient fault propagation path, whether the pupative element fails by 709 in tension or shear. Elements that do not fail by one of these criteria do not alter W_{ext} , so the 710 efficiency of the system is unchanged. Thus, by utilizing these failure criteria along 711 pupative elements, the principle of work minimization is applied and theories governing 712 fault mechanics are honored. This approach has the advantage of accommodating both in-713 line growth and growth at angles to pre-existing faults without invoking separate tensile 714 and shear failure criteria that require parallel analyses. The work minimization approach 715 also proves to be a viable alternative to using Coulomb theory, which empirically captures 716 the physics of faulting, but is limited in its ability to predict propagation at fault tips. The 717 main limitation is in generation of two equally viable potential failure planes. While one of 718 the failure planes may match the observations of fault growth, Coulomb theory does not 719 provide a means for objectively picking one plane over the other. While Coulomb theory is 720 remarkably helpful for predicting failure of pre-existing structures, or comparing failure 721 potential between multiple fault surfaces, it may not be so for predicting fault propagation 722 and evolution.

723 We see two main applications of the principle of work minimization to analyses of 724 geologic structure. The first is to use the full energy budget, or its individual components, 725 to gain insight into processes of deformation. For example, recent work by Savage et al. 726 (2014) constrained the energy required for fault frictional heating by estimating 727 temperatures from biomarkers near an exhumed fault. Work against friction also was 728 estimated recently by Fulton et al. (2013) from in-situ measurements of a fault that failed 729 in the 2011 Tohoku earthquake. Similarly, the work budget within accretionary systems 730 has shed light on the tradeoffs of underthrusting and accretionary phases in wedge 731 development (e.g. Gutscher et al., 1998; Hardy et al., 1998; Burbridge and Braun, 2002; Del 732 Castello and Cooke, 2007; Mary et al., 2013). Other studies that could benefit from similar 733 analyses include the partitioning of off-fault deformation between different faults, the 734 trade-off between various processes and uplift against gravity, and how different faulting 735 scenarios affect estimates of the energy available for seismic shaking.

The second application of work minimization to structural geology is the predictionof fault propagation paths. This analysis has been done by comparing the total or external

- 738 work of different interpretions of fault configurations along the southern San Andreas fault
- over the past 500,000 years (Cooke and Dair, 2011). Alternatively, work minimization can
- be used in conjunction with numerical modeling tools to predict propagation paths over
- 741 multiple growth cycles, as presented here. The development of software to automate this
- 742 process would provide powerful tools for predicting fault growth.

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930 Figures

- 931 Figure 1. Schematic of fault deformation shows the tectonic work (*W*_{ext}), internal work of
- 932 deformation (W_{int}), work of uplift against gravity (W_{grav}), the work against frictional slip
- 933 (W_{fric}) , the energy required to grow a fault (W_{prop}) and the radiated seismic energy (W_{seis}) .
- Taken from Cooke and Murphy, 2004.
- 935 Figure 2. Schematic of the partitioning of frictional heating, the energy required for fault
- growth and energy available for seismic shaking. Work is done when fault slips and the
- 937 shear stress drops from the initial value, τ_0 , to the value at stable sliding, τ . One can
- 938 delineate the associated work terms in different ways. In most studies, *W*_{fric} is defined as
- 939 the energy due to slip under sliding shear stress, consequently *W*_{fric} depends on absolute
- 940 stress. The two energy terms that are consumed by fault growth (W_{seis} and W_{fric}) depend on
- 941 the shear stress drop during the slip event. Prior to the slip reaching the slip weakening
- 942 distance, *L*, work is primary consumed by the generation of new fault surface area (W_{prop})
- 943 and *W*_{fric}. Slip beyond the slip weakening distance produces *W*_{fric} and seismic radiated
- 944 energy (*W*_{seis}). While the sketch in (A) shows this relationship for a single increment of
- loading the hypothetical sketch in (B) shows how W_{fric} , W_{prop} and W_{seis} are calculated over
- 946 multiple loading increments, which are typically implemented within numerical models.
- Figure 3. Minimum viscous dissipation demonstrates that of the possible ocean ridge (red)and transform (blue) intersections, several of which are show here, geometry (C) is the
- 949 most efficient. Modified from Lachenbruch and Thompson, 1972.
- 950 Figure 4. A) Strain energy density (SED) and B) Coulomb stress for the interpreted Puente
- Hills fault system prior to the growth of the Coyote Hills fault (CH) in southern California.
- 952 The region of high SED envelops the location of the Coyote Hills fault. If we use high SED
- 953 and extreme values of Coulomb stress to predict the development of the next fault, the SED
- 954 pattern gives a better correlation with the location of the observed fault. The Coulomb
- 955 stress is less conclusive than the SED and suggests that the next faults may develop either
- 956 northeast or southwest of the Whittier fault. Taken from Olson and Cooke, 2005.
- Figure 5. (A) Limit analysis can be used to predict the sequence of thrusting. (B) Thepredicted evolution of the tectonic force shows drops in force with the growth of new

959 faults. Taken from Cubas et al. (2008).

960 Figure 6. DelCastello and Cooke (2007) explored the evolving work budget of forethrust 961 growth within the analog experiments of Adam (2005). (A) The numerical models simulate 962 lengthening of the thrust sheet prior to forethrust development. (B) With lengthening of 963 the thrust sheet length, W_{fric} increases as the fault normal compression increases due to the 964 thickening wedge. Consequently, the fault system becomes increasingly inefficient until a 965 new forethrust grows. This fault growth reduces the W_{ext} by decreasing W_{fric} along the fault 966 by an amount greater than the increase in *W*_{int} associated with the addition of the new fault. (C) If the forethrust were not to grow, the W_{ext} would continue to increase. The timing of 967 968 fault growth depends on the work required to grow the new fault surface. The new fault 969 grows at 10.5 cm thrust fault length rather than sooner in the experiment because the total 970 energy reduction due to including the fault within the model (prefault W_{ext} – post fault W_{ext}) exceeds the energy consumed in the growth of the fault (W_{prop} and W_{seis}). Taken from 971 972 DelCastello and Cooke, 2007.

Figure 7. Two models based on laboratory experiments on faults within gypsum slabs by Bobet and Einstein (1998). (A) Set-up for the numerical experiment of a 12.7 mm joint in a 160 x 80 mm slab of gypsum. The block is subject to 27.1e⁻⁶ m of uniaxial extension (u_v). E is Young's modulus and v is Poisson's ratio. (B) Set-up for the numerical experiment on a 12.7 mm fault within gypsum that dips 60° from the vertical compressive stress, $\sigma_v = -29$ MPa. A confining pressure of $\sigma_h = -2.5$ MPa also is applied. Box shows location of (C). Box at

979 fault tip is the location of the inset, showing pupative elements radial to the fault tip.

Figure 8. ΔW_{ext} per pupative element area from before growth to after growth of (A) the

981 joint and (B) the fault models. Red markers are pupative element orientations where

neither the tensile nor the shear failure criterion is met, so no growth can occur and there

983 is no change in work. Green markers are orientations where the tensile strength is

984 exceeded and failure can occur by opening. Blue markers are orientations where shear

failure can occur. Tensile strength, T, is 3.2 MPa. Shear strength, S_o, is 78.7 MPa.

986 Figure 9. Work budgets for joint and fault propagation by one pupative element in the

987 orientation that maximizes ΔW_{ext} . The energy to grow faults ($W_{prop} + W_{seis}$) comes from

- 988 ΔW_{int} . Growth elements for Fault 1 are not slip-weakening and do not experience a drop in
- 989 strength during sliding. The inherent friction, $\mu_0 = 0.3$, is equal to the dynamic friction, μ_d .
- Growth elements for Fault 2 are slip-weakening, with $\mu_o = 0.6$, $\mu_d = 0.3$, and L = 4x10⁻⁶ m.
- Growth elements for Fault 3 also are slip weakening, with $\mu_0 = 0.6$, $\mu_d = 0.3$, and L = 1x10⁻⁶
- m. The most efficient propagation path for both the modeled joint and for all three faults is
- 993 collinear at 180°.
- Figure 10. (A) Planes carrying maximum tensile stress are perpendicular to the maximum
- 995 tensile stress, σ_1 (top, green), while the two planes carrying the maximum Coulomb shear
- 996 stress are oriented at angle φ to the maximum compressive strength, σ_3 (bottom, red). (B)
- 997 Contours of the maximum tensile stress. (C) Contours of the maximum Coulomb shear
- 998 stress, σ_c , on optimally oriented planes with $\mu_o = 0.3$. (D) Tensile and Coulomb failure
- 999 planes at 5 points located at one half the element length from the fault tip (0.0635 mm).
- 1000 Tensile strength, T, is 3.2 MPa. Shear strength, S_o, is 78.7 MPa. Strength must be exceeded
- 1001 for failure to occur (solid lines). Tensile and shear failure can occur at a variety of points
- around the fault tip, but this approach does not permit prediction of the favored
- 1003 propagation path for the fault.



Figure 2























Scenario	μs	μ_d	L	W _{prop}	Wseis	ΔW_{fric}	$W_{prop} + W_{seis} + \Delta W_{fric}$	ΔWint	ΔW _{ext}
			(m)	(J/m²)	(J/m²)	(J/m²)	(J/m²)	(J/m²)	(J/m²)
Joint	0	0	0	3.62	0.00	0.00	3.62	3.63	3.63
Fault 1	0.3	0.3	0	260.77	0.00	79.78	340.54	254.05	333.83
Fault 2	0.6	0.3	1e⁻6	158.60	102.14	79.45	340.19	254.38	333.83
Fault 3	0.6	0.3	4e ⁻⁶	39.70	221.11	79.45	340.26	254.38	333.83

Table 1: Properties and work budgets for growing joint and fault models