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4 5 **Rapid Fault Leakage Screening for CO₂** 6 **Storage Under Uncertainty:** 7 **Application to a Faulted Miocene** 8 **Aquifer in the Malay Basin**

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34 **Rapid Fault Leakage Screening for CO₂ Storage Under**
35 **Uncertainty: Application to a Faulted Miocene Aquifer in**
36 **the Malay Basin**

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53 **Abstract**

54 Geological CO₂ storage is critical for mitigating climate change, yet the risk of leakage
55 along faults poses significant challenges. Faults function as structural traps or
56 preferential CO₂ migration pathways depending on their sealing capacity. This study
57 applies a computationally efficient rapid fault leakage analysis tool that combines a
58 vertically integrated reservoir model with an upscaled fault leakage function based on
59 steady-state flow approximations, conceptualising faults as zones of increased vertical
60 and reduced horizontal permeability. The methodology is applied to faulted Miocene
61 clastic aquifers in the Malay Basin, representing a complex, structurally heterogeneous
62 setting typical of mature Southeast Asian storage prospects. The analysis evaluates 600
63 well locations to identify optimal injection sites, incorporating both deterministic
64 scenarios (fully sealing versus leaking faults) and comprehensive uncertainty
65 quantification across fault property heterogeneity. Comparison of sealing and leaking
66 scenarios establishes physical controls governing leakage behaviour, while uncertainty
67 quantification using 10,000 Monte Carlo realisations across 71 parameters (fault-
68 specific permeability, capillary entry pressure, and transmissibility for 23 faults, plus
69 reservoir properties) identifies the dominant controls on leakage outcomes. First,
70 leakage sensitivity concentrates on a small fault subset: only 5 of 23 faults exhibit
71 Spearman correlation with outcomes, with a single fault dominating system response.
72 Second, k-medoids clustering identifies three distinct regimes, with high-leakage
73 scenarios comprising 44% of realisations, the modal outcome rather than a rare event
74 under current parameter uncertainty. Third, cluster membership is predominantly
75 controlled by Fault 23 permeability, while Fault 1 transmissibility controls lateral access
76 to the northern fault array, a cascade pathway that demonstrates the value of multi-fault
77 uncertainty quantification. The findings demonstrate that targeted characterisation of 5-

78 7 critical faults can effectively collapse predictive uncertainty, enabling transition from
79 high-uncertainty screening to confident operational planning. The tool's computational
80 efficiency enables rapid assessment supporting early-stage site screening, risk
81 assessment, and targeted data acquisition, thereby accelerating the identification and
82 prioritisation of secure CO₂ storage sites.

83

84 **Keywords:** CO₂ storage, fault leakage, vertical equilibrium modelling, cluster analysis,
85 faulted aquifers, uncertainty quantification

86 1 Introduction

87 Geological CO₂ storage is indispensable for mitigating climate change and achieving
88 global decarbonization targets (Krevor et al. 2023). Malaysia's National Energy Transition
89 Roadmap exemplifies this urgency, targeting CO₂ storage capacities of 15 million tonnes
90 per annum (Mtpa) by 2030 and 80 Mtpa by 2050 (Fernandez et al., 2024). Meeting these
91 targets requires systematic evaluation of storage prospects across diverse geological
92 settings to ensure long-term safe containment within these formations. Leakage
93 compromises mitigation efforts, poses environmental and health risks, and can
94 undermine public confidence (Bielicki et al., 2015; Ashworth et al., 2015; Jones et al.,
95 2017). Regulatory frameworks therefore require comprehensive geological
96 characterisation and rigorous site assessment to identify and mitigate leakage risks
97 (Dixon et al., 2015).

98 Faults play a critical role in storage viability, functioning either as structural traps that
99 enhance containment or as preferential pathways for upward CO₂ leakage when their
100 sealing capacity is compromised (Caine et al., 1996; Dixon et al., 2015). Fault zone
101 architecture produces a strong permeability contrast between fault zone components:
102 fault cores exhibit permeabilities several orders of magnitude below the surrounding
103 protolith, while fractured damage zones can be 1–3 orders of magnitude more permeable
104 than the host rock, creating preferential pathways for along-fault fluid migration (Bense
105 and Pearson, 2006; Faulkner et al., 2010; Bense et al., 2013; Scibek, 2020). Damage zone
106 width scales with fault displacement, and permeability varies systematically with
107 fracture density and connectivity across the fault zone (Childs et al., 2009; Faulkner et
108 al., 2011). The capillary sealing capacity of fault damage zones is governed by fracture
109 aperture distributions and clay content rather than the intact matrix entry pressures
110 relevant to unfaulted caprocks, with effective values spanning several orders of
111 magnitude depending on fracture fill mineralogy and connectivity (Guiltinan et al., 2018;
112 Phillips et al., 2020).

113 Hydraulic properties of fault zones are poorly constrained at early screening stages and
114 span multiple orders of magnitude even within the same geological setting (Miocic et al.,
115 2019b; Scibek, 2020). This parameter uncertainty translates directly into outcome
116 uncertainty by substantially altering predicted storage capacity in fault-bounded
117 reservoirs. Recent ensemble-based approaches have addressed fault leakage
118 uncertainty through stochastic upscaling with copula representation (Pettersson et al.,
119 2025), deep-learning surrogates for coupled flow-geomechanics (Lu et al., 2025), and
120 probabilistic transmissibility models (Salo-Salgado et al., 2025), but each of these
121 studies characterises a single fault zone in isolation. In storage settings with many faults,
122 identifying which fault properties drive leakage outcomes, and therefore which

123 measurements would most efficiently reduce predictive uncertainty, requires systematic
124 ensemble analysis across the full fault population.

125 Traditional reservoir simulation approaches using fine-scale coupled models are
126 computationally prohibitive for evaluating multiple sites or conducting comprehensive
127 uncertainty quantification with thousands of parameter combinations (Birkholzer et al.,
128 2012; Ringrose & Meckel, 2019). These combined data and computational limitations
129 constrain early-stage decision making. Rapid assessment tools can address these
130 limitations by accelerating CO₂ storage deployment through reduced early-stage project
131 risks and costs while maintaining sufficient accuracy for screening purposes.

132 In this study, we address this early stage in the exploration funnel and build on the rapid
133 fault leakage framework developed by Ramachandran et al. (2026). Our previous work
134 combines a vertically integrated reservoir model with an upscaled fault leakage function
135 based on steady-state flow approximations. Here, faults are conceptualised as zones of
136 increased vertical permeability in the caprock and reduced horizontal permeability in the
137 reservoir. We integrate this rapid fault-leakage modelling framework with previous
138 studies focused on geological screening and CO₂ storage site selection for the Malay
139 Basin characterisation (de Jonge-Anderson et al., 2024a, 2024b, 2025a). It is applied to a
140 predefined optimal storage zone in the Lower Miocene Group J aquifer, a sand-rich unit
141 with high net-to-gross, elevated porosity and permeability, and complex faulting typical
142 of mature Southeast Asian basins (de Jonge-Anderson et al., 2025a). The analysis
143 combines deterministic sealing and leaking scenarios with a 10,000-realisation Monte
144 Carlo uncertainty quantification across 71 parameters (fault permeability, capillary entry
145 pressure, and transmissibility multiplier for each of 23 faults, plus reservoir porosity and
146 permeability exponents). Spearman rank correlation quantifies the sensitivity of leakage
147 outcomes to each parameter, and k-medoids clustering partitions the ensemble into
148 distinct behavioural regimes with medoid scenarios available for detailed follow-up
149 simulation. Conditional cluster probability analysis then identifies the parameter values
150 at which regime membership shifts, translating the sensitivity ranking into useful site
151 characterisation targets.

152 2 Geological Background

153 The Malay Basin is a 70,000 km² extensional basin located off the east coast of Peninsular
154 Malaysia (Figure 1a). It has a long history of hydrocarbon exploration and production, with
155 14.8 billion barrels of oil equivalent produced from multiple reservoirs spanning its thick
156 (up to 13 km) Cenozoic sedimentary fill (Madon, 2021). These characteristics have led to
157 its identification as an attractive basin for CO₂ geological storage, with initial studies
158 confirming gigatonnes of storage capacity across multiple aquifers and depleted
159 hydrocarbon intervals (Hasbollah et al., 2020; Zhang & Lau, 2022; de Jonge-Anderson et
160 al., 2025a).

161 This study specifically focuses on the Group J stratigraphic interval, selected based on its
162 identification as a viable aquifer storage target in the regional GCS screening of de Jonge-
163 Anderson et al. (2025a). It was deposited during the lower Miocene in shallow marine
164 facies and consists of high-quality sandstone reservoirs. The aquifer represents one of
165 the higher net-to-gross (NTG) sequences in the basin (de Jonge-Anderson et al., 2025a).
166 Burial diagenesis significantly modifies reservoir properties, with porosity declining
167 rapidly with depth, particularly within the uppermost 2000 m of burial. However, within
168 optimal zones (< 3000 m burial depth), a median porosity of 18 % is maintained.

169 The Malay Basin is affected by pronounced faulting, particularly evident along the
 170 western margin hinge zone and central basin areas where north-south trending faults
 171 create extensive horst and graben systems that exert significant control over the depth
 172 structure and compartmentalisation of the basin fill (de Jonge-Anderson et al., 2024b).
 173 An *en-echelon* faulting pattern of overlapping curved faults affects the north of the basin.
 174 As a mature hydrocarbon province, the basin's faults have historically acted as effective
 175 seals for hydrocarbons over geological timescales, demonstrating inherent sealing
 176 capacity. However, seismic evidence indicates fault leakage in some locations (de Jonge-
 177 Anderson et al., 2024b), and deep-seated faults likely act as migration pathways for
 178 naturally occurring CO₂ (Madon et al., 2004). This study assesses the risk of CO₂ leakage
 179 along faults under injection conditions in this mature basin.
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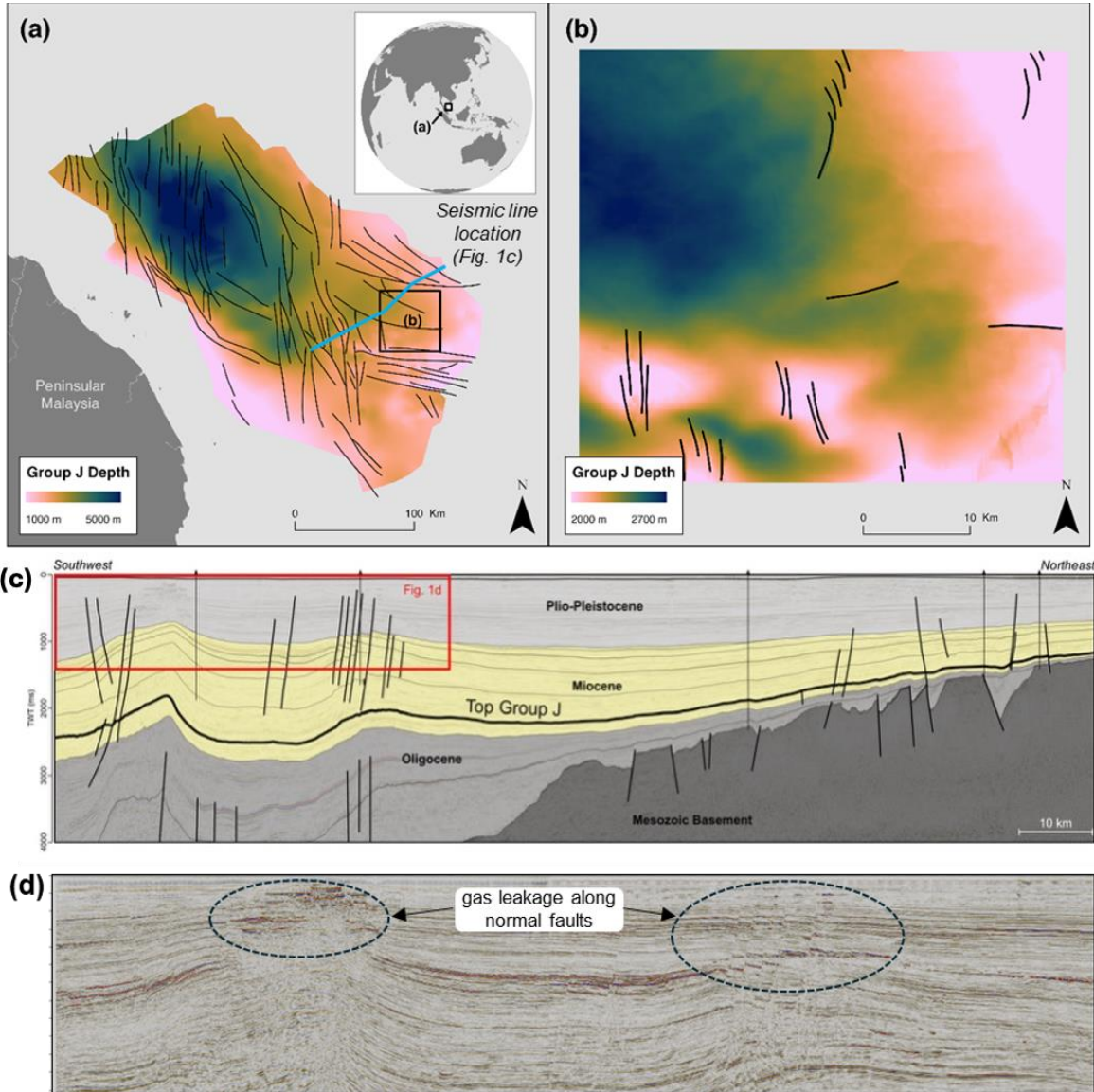


Figure 1 – Location maps of the Malay Basin and study area. (a) Regional depth map of the top Group J stratigraphic surface, constructed from horizon data and regional fault maps of de Jonge-Anderson (2025b) and PETRONAS (2022). (b) Local map of the study area at the eastern basin margin within the predefined optimal CO₂ storage zone; faults mapped from 3D seismic data as part of this study. (c) Regional structural cross-section adapted from de Jonge-Anderson et al. (2024b). (d) Seismic section showing acoustic anomalies consistent with vertical fluid migration.

181 3 Methods

182 3.1 3D Seismic Mapping

183 The area of interest lies at the southeast margin of the Malay Basin (Figure 1b). We used
184 time-domain 3D seismic datasets to map a highly reflective and laterally continuous
185 seismic reflector near the top of Group J. The horizon was gridded at 500 m cell resolution,
186 then depth-converted using a generalized time-depth relationship from regional
187 checkshot data (de Jonge-Anderson et al., 2024b). Faults were interpreted from visible
188 offsets of the mapped horizon. The area slopes gently downdip to the northwest from
189 approximately 2000 m in the east to around 2700 m in the west. In the south, a series of
190 west-east oriented anticlines locally uplift Group J to approximately 2000 m. We
191 constructed a simple 3D geological model from this mapped surface. The model top
192 corresponds to the Group J surface, and the base was defined by duplicating this horizon
193 and shifting it down by 100 m, representing the average Group J thickness in this area.

194 3.2 Reservoir Modelling

195 The reservoir modelling is based on fundamental equations governing fluid flow in porous
196 media. The 3D mass conservation equation for two immiscible and incompressible fluid
197 phases α , CO₂ ($\alpha = g$) and brine ($\alpha = w$) turn into a volume conservation equation

$$\frac{\partial(\phi s_{\alpha})}{\partial t} + \nabla \cdot u_{\alpha} = q_{\alpha}, \quad (1)$$

198 where s_{α} is the saturation of phase α , ϕ is the porosity, u_{α} is the Darcy velocity of phase
199 α , and q_{α} is a source/sink term in units of volume of phase α per time. The porous medium
200 is assumed rigid and isothermal. The volume balance is established by

$$s_g + s_w = 1. \quad (2)$$

201 The Darcy velocity is given by

$$u_{\alpha} = -\frac{k_{r\alpha}}{\mu_{\alpha}} k(\nabla p_{\alpha} - \rho_{\alpha} \mathbf{g}), \quad (3)$$

202 where $k_{r\alpha}$ is the relative permeability, μ_{α} is the viscosity, k is the permeability tensor, ρ_{α}
203 is the density, p_{α} is the fluid pressure (all for phase α), and \mathbf{g} is the gravity acceleration
204 vector. The phase pressures are related by the capillary pressure (p_c) function

$$p_c = p_g - p_w. \quad (4)$$

205 To obtain the fluid saturations and pressures by solving this system, three conditions
206 must be met: firstly, specific functions must represent relative permeability and capillary
207 pressure (Brooks-Corey functions used here); secondly, initial conditions for pressure
208 and saturation throughout the reservoir domain must be established; and thirdly,
209 appropriate boundary conditions must be defined.

210 3.3 Vertical Equilibrium Modelling

211 Vertical Equilibrium (VE) models exploit two fundamental characteristics of CO₂ storage
212 systems: strong buoyancy contrast between CO₂ and brine, and large aspect ratio where
213 lateral reservoir dimensions (hundreds of meters to kilometres) vastly exceed vertical
214 dimensions (meters to tens of meters) (Huppert & Woods, 1995; Yortsos, 1995). These
215 characteristics promote rapid vertical CO₂ segregation relative to lateral migration,
216 establishing vertical hydrostatic equilibrium. Consequently, CO₂ accumulates as thin
217 layers beneath low-permeability seals, and vertical pressure and saturation distributions

218 can be approximated by balancing buoyancy and capillary forces (Nordbotten & Celia,
219 2011).

220 VE models reduce problem dimensionality by vertically averaging the 3D conservation
221 equations and Darcy's law, transforming the system into 2D governing equations solved
222 on the lateral plane (Andersen et al., 2015; Nilsen et al., 2015). The approach preserves
223 the standard multiphase Darcy equation structure, enabling implementation in
224 conventional reservoir simulators and extension to incorporate additional physical
225 phenomena including capillary hysteresis (Doster et al., 2013), residual trapping (Nilsen
226 et al., 2016), and geomechanical coupling (Bjornara et al., 2016; Andersen et al., 2016,
227 2017). Post-simulation reconstruction recovers vertical pressure and saturation
228 distributions from vertically integrated variables when detailed 3D fields are required.
229 Benchmark studies comparing VE simulations against fully three-dimensional models for
230 field-scale problems, including the Sleipner injection project, demonstrate comparable
231 accuracy with orders-of-magnitude computational speedup (Nilsen et al., 2011, 2015,
232 2016; Bandilla et al., 2014; Moyner & Nilsen, 2019). We use a VE model implemented
233 within the open-source software package MRST-co2lab, a module within the MATLAB
234 Reservoir Simulation Toolkit (MRST) (Andersen, 2017; Lie, 2019). For comprehensive
235 derivations and detailed treatment of VE assumptions and limitations, readers are
236 referred to Nordbotten & Celia (2011). The simulation outputs are further used to identify
237 an optimal injection well location through sequential filtering on overpressure, boundary
238 connectivity, and CO₂ retention, as described in Section 4.1.3.

239 3.4 Upscaled Fault Leakage Function

240 The fault leakage modelling approach employs the analytical framework of
241 Ramachandran et al. (2026), which estimates leakage rates using steady-state
242 approximations. The fault is conceptualised as a permeable damage zone connecting the
243 storage reservoir to an overlying aquifer (Figure 2). Following established understanding
244 of fault zone architecture (Faulkner et al., 2010), the fault comprises a low-conductivity
245 core surrounded by high-conductivity damage zones. The fault core acts as a barrier to
246 across-fault flow within the reservoir, while the damage zones, containing fracture
247 networks and deformation bands, permit upward fluid migration. For computational
248 efficiency, the fault core is treated as impermeable (no-flow boundary), and the damage
249 zone is modelled as an equivalent porous medium with representative bulk properties.
250 This simplified representation provides a worst-case scenario for vertical leakage
251 assessment. Hereafter, this damage zone is referred to as "the fault zone" or "the fault".
252 Prior to CO₂ reaching the fault base, reservoir overpressure can drive brine leakage along
253 the fault zone. The vertical brine leakage flux Q_{wf} is given as,

254

$$Q_{wf} = \begin{cases} 0, & h_g > 0 \\ \frac{A_f k_f (P_w - P_{w0})}{\mu_w L_c}, & h_g = 0 \end{cases} \quad (5)$$

where A_f is the area of the fault perpendicular to flow, k_f is the vertical fault permeability, μ_w is brine viscosity, h_g is the CO₂ column height in the reservoir block connected to the fault block, P_w is the brine pressure in the reservoir, P_{w0} is the initial brine pressure in the reservoir, and L_c is the caprock thickness (i.e., the fault length connecting the storage reservoir to an overlying aquifer, assuming a vertical fault geometry for simplicity). Brine leakage occurs when the reservoir pressure exceeds the

initial pressure at the fault base, providing a pressure relief mechanism that reduces reservoir overpressure before it ceases once the CO₂ plume makes contacts. We assume that once the CO₂ plume contacts the fault base, the CO₂ plume forms a sufficiently thick layer that excludes the aqueous phase entry at low leakage rates (Kang et al., 2014).

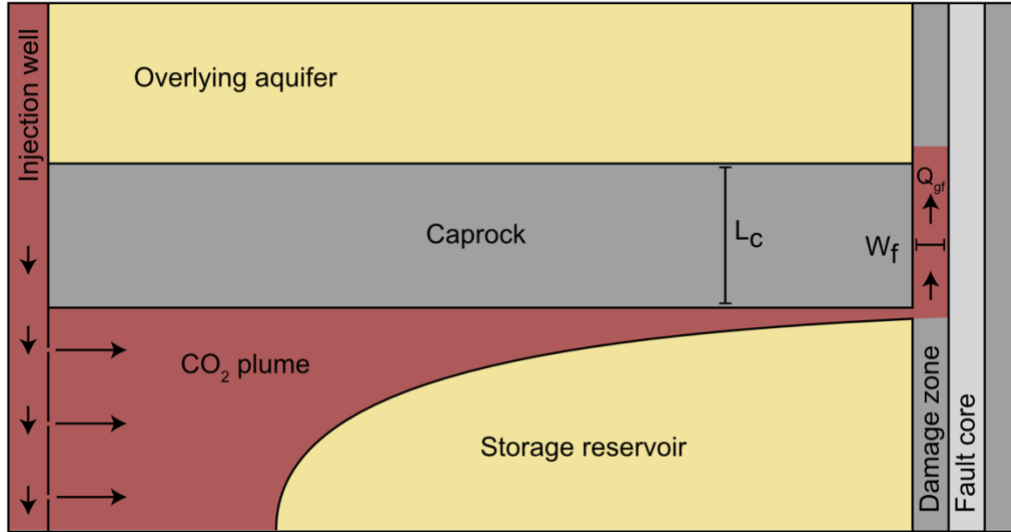


Figure 2 - Schematic vertical cross-section of a faulted storage reservoir showing fault architecture and leakage pathway. CO₂ leaks along the fault damage zone once it reaches the fault base within the reservoir. The fault core acts as a barrier to horizontal flow, while the damage zone permits vertical migration to a shallower overlying aquifer. Adapted from Ramachandran et al. (2026).

255

256 The along-fault leakage calculation adapts the analytical solution for flow through
 257 fissures between aquifers (Neufeld et al., 2009) to account for reservoir overpressure
 258 from injection and fault sealing capacity (Ramachandran et al., 2026). Fault seal capacity
 259 is controlled by capillary entry pressure (p_e), representing the threshold pressure
 260 required for CO₂ to invade the fault zone. The driving potential (ψ) is modelled as

$$\psi = \Delta\rho g h_g + (P_w - P_{w0}) - p_e, \quad (6)$$

261 where $\Delta\rho = \rho_w - \rho_g$ is the density difference between brine and CO₂, and p_e is the fault
 262 capillary entry pressure. The first term represents buoyancy-driven flow, while the second
 263 term captures pressure-driven flow from injection-induced overpressure. CO₂ enters the
 264 fault base when $\psi > 0$, representing the capillary entry condition. Once entry is achieved,
 265 along-fault migration through the damage zone is governed by Darcy flow (Eq. 7). In the
 266 steady-state single-phase formulation adopted here, entry at the fault base and
 267 breakthrough at the fault top are treated as simultaneous: the $\Delta\rho g L_c$ term in Eq. 7
 268 represents the buoyancy head that builds as CO₂ traverses the fault zone of thickness L_c ,
 269 such that the breakthrough pressure at the fault top is $p_e + \Delta\rho g L_c$, consistent with
 270 laboratory measurements of capillary seal capacity (Busch & Amann-Hildenbrand, 2013;
 271 Espinoza & Santamarina, 2017). Higher capillary entry pressures delay leakage onset and
 272 reduce cumulative leakage volumes. The vertical gas leakage flux Q_{gf} along the fault is
 273 given as

$$Q_{gf} = \begin{cases} 0, & \psi \leq 0 \\ \frac{A_f k_f (\psi + \Delta\rho g L_c)}{\mu_g L_c}, & \psi > 0 \end{cases}, \quad (7)$$

274 where μ_g is the CO₂ viscosity. This approach relaxes the VE assumption at grid blocks
275 connected to the fault, allowing non-zero vertical flow. However, for steady-state single-
276 phase flow, fault leakage has a minimal effect on reservoir predictions (Kang et al., 2014).
277 The fault leakage function is implemented as a source/sink term in the vertically
278 integrated mass balance equations, avoiding explicit 3D fault discretisation. Grid cells
279 containing faults have their across-fault permeability reduced using transmissibility
280 multipliers or set to zero (representing the impermeable core) and incorporate the fault
281 leakage function with specified properties. The steady-state single-phase assumption
282 neglects transient accumulation within the fault. Although considerable, the impact of
283 stress-sensitive permeability or fault reactivation is neglected here (Rutqvist, 2012;
284 Rinaldi et al., 2014; Jha and Juanes, 2014; Vilarrasa et al., 2017; Zhao and Jha, 2019;
285 Meguerdijian and Jha, 2021). These simplifications are acceptable for early-stage
286 screening where the objective is bounding plausible leakage ranges rather than precise
287 prediction. Full details on numerical implementation, verification, and sensitivity
288 analyses are provided in Ramachandran et al. (2026).

289 3.5 Data Analysis and Clustering

290 Fault properties exhibit substantial uncertainty, especially during early-stage storage site
291 assessment due to limited subsurface data. To quantify this uncertainty and identify
292 parameter combinations controlling leakage behaviour, we employ Monte Carlo
293 simulation with Latin hypercube sampling (McKay et al., 1979) combined with Spearman
294 rank correlation analysis and k-medoids clustering. The uncertainty analysis samples 71
295 parameters independently: fault permeability, fault capillary entry pressure, and fault
296 transmissibility multiplier for each of the 23 faults, plus reservoir porosity and
297 permeability exponents. Fault properties follow log-uniform distributions spanning
298 multiple orders of magnitude: k_f from 10^{-6} to 10^{-2} mD, p_e from 0.01 to 10 bar, and fault
299 transmissibility multiplier from 10^{-6} to 10^{-1} . The exponents governing the porosity-depth
300 and permeability-porosity relationships (Eq. 8 and 9) follow truncated normal
301 distributions reflecting petrophysical relationships calibrated to basin data (de Jonge-
302 Anderson et al., 2025a). Each fault is sampled independently to quantify fault-specific
303 heterogeneity effects rather than assuming uniform properties across the fault system.
304 Latin hypercube sampling ensures efficient parameter space coverage with fewer
305 samples than random Monte Carlo methods while maintaining computational
306 tractability. Spearman rank correlation analysis quantifies individual parameter
307 sensitivity, selected for its robustness to non-linear relationships and log-distributed
308 outcomes (Saltelli et al., 2007). For each parameter, the Spearman correlation coefficient
309 between parameter value and total fault leakage identifies which faults and properties
310 dominate system response.

311 K-medoids clustering partitions realisations into distinct behavioural regimes based on
312 their positions in the parameter space. Unlike k-means clustering, which computes
313 abstract centroids, k-medoids selects actual data points (medoids) as cluster
314 representatives (Park & Jun, 2009). This distinction is critical for subsurface uncertainty
315 analysis: medoids correspond to actual parameter combinations that can be extracted
316 for detailed verification using computationally intensive 3D coupled simulations (Scheidt
317 & Caers, 2009; Mahjour & Faroughi, 2023). Alternative approaches such as distance-
318 based global sensitivity analysis (DGSA) identify influential parameter regions but do not
319 produce specific testable scenarios. Several methods exist for determining the optimal
320 number of clusters, including the elbow method, silhouette analysis, and the gap statistic

321 (Freites et al., 2023); the elbow method is used here, identifying the inflection point in
322 within-cluster sum of squares versus cluster count. Unlike traditional sensitivity analysis
323 that examines individual parameter influence, clustering reveals which parameter
324 combinations define distinct leakage regimes, each encompassing a range of outcomes
325 represented by a medoid scenario. This is a critical distinction because real-world
326 leakage results from multi-parameter interactions rather than single-factor controls
327 (Rinaldi et al., 2014; Gasda et al., 2022; Mahjour & Faroughi, 2023). The clustering
328 identifies endmember scenarios representing best-case (minimal leakage) and worst-
329 case (maximum leakage) outcomes, which can be extracted for detailed 3D numerical
330 simulations without requiring full computational evaluation of the entire parameter
331 space. Prior to clustering, all parameters are normalised using robust median-IQR
332 (interquartile range) scaling: log-uniform fault parameters and total fault leakage are
333 \log_{10} -transformed before scaling, whilst the normally distributed reservoir exponents are
334 scaled on their original scale, ensuring comparable contributions to the Euclidean
335 distance metric. Total fault leakage is included as an additional variable with a weight of
336 2, ensuring that cluster boundaries reflect leakage behaviour rather than parameter-
337 space proximity alone.

338 To characterise the functional relationship between individual parameters and leakage
339 regime membership, we compute the conditional cluster probability $P(\text{cluster}|\mathbf{x})$ for each
340 of the top 15 parameters ranked by spearman correlation coefficient. For each discrete
341 bin along the parameter's sampled range, conditional cluster probability is estimated as
342 the fraction of all realisations falling within that bin that were assigned to a given cluster.
343 Bin edges are defined uniformly across the full ensemble range of each parameter using
344 20 equal-width intervals, with a minimum occupancy threshold of 10 realisations per bin
345 applied to suppress unreliable estimates at sparse distribution tails. The unconditional
346 cluster fractions, defined as the fraction of all realisations assigned to each cluster
347 irrespective of any single parameter value, serve as the null baseline. A parameter with
348 no discriminating power produces conditional probability curves that coincide with these
349 baselines across its full sampled range.

350 Together, these methods translate a high-dimensional uncertainty quantification into
351 actionable site characterisation targets: rather than requiring comprehensive
352 characterisation across all 71 parameters, the sensitivity ranking and conditional
353 probability analysis identify the small subset of faults whose properties are most critical
354 to storage outcomes. These are the few faults, among the many, that warrant targeted
355 characterisation. This targeted approach is particularly valuable at early-stage screening,
356 where full-physics 3D coupled simulations are computationally prohibitive across the full
357 parameter space.

358 4 Results

359 4.1 Well Location Identification Under Fault Uncertainty

360 4.1.1 Model Parameters

361 The model is applied to a section of Group J in the Malay Basin, as illustrated in Figure 3.
362 Key geological and operational parameters are listed in Table 1, including a seafloor
363 depth of 70 m below sea level, an average top reservoir depth of 2300 m below the
364 seafloor, a geothermal gradient of $50\text{ }^{\circ}\text{C.km}^{-1}$, and a seafloor temperature of $24\text{ }^{\circ}\text{C}$ (Madon
365 & Jong, 2021; de Jonge-Anderson et al., 2025a). Rock compressibility is applied to the
366 reservoir formation only; the value adopted here is from de Jonge-Anderson et al. (2024a)

367 for a neighbouring field within the Malay Basin. The reservoir model has a grid resolution
 368 of 500 m x 500 m (Figure 3). The porosity (ϕ) and permeability (k) across the grid are
 369 obtained from de Jonge-Anderson et al., (2025a), where porosity is calculated as a
 370 function of depth (z) and permeability is derived from porosity (Eq. 8 and 9; Figure 4).

$$\phi = 0.45e^{(-a_1z)} \quad (8)$$

$$k = 700236\phi^{b_1} \quad (9)$$

371 where a_1 (10th/median/90th percentile = $3.83 \times 10^{-4}/5.34 \times 10^{-4}/6.84 \times 10^{-4}$) follows a
 372 truncated normal distribution (mean = 5.34×10^{-4} and standard deviation = 2.15×10^{-5}),
 373 and b_1 (10th/median/90th percentile = $3.47/4.05/4.63$) follows a truncated normal
 374 distribution (mean = 4.05 and standard deviation = 8.22×10^{-2}). Section 4.1 uses median
 375 values for deterministic fault location, while Section 4.2 employs the distributions for
 376 uncertainty quantification. Permeability (k) is in millidarcies (mD). Fault width is fixed at
 377 5 m, representative of faults with 10 - 100 m displacement in clastic settings (Childs et
 378 al., 2009; Torabi & Berg, 2011). Because damage-zone width enters the upscaled leakage
 379 function as a product with fault permeability (Eq. 7), uncertainty in width is implicitly
 380 subsumed within the sampled fault permeability.

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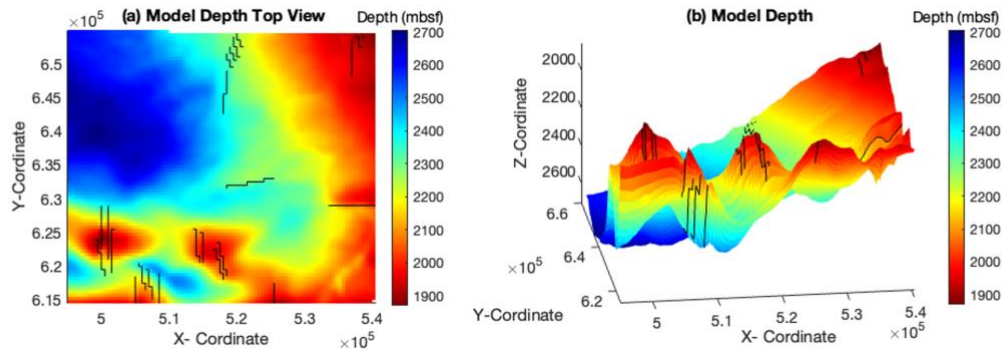


Figure 3 - Depth structure of the Group J top surface. (a) 2D map showing the spatial variation in depth. The colour scale represents depth from deep (blue) to shallower (red). Black lines show the 24 mapped faults. (b) 3D surface plot showing the topographical distribution. The z-axis represents the depth. Both visualisations reveal significant depth variations, with shallower regions in the central and northeastern portions and a structural depression in the southwest.

382

383 We explicitly define 23 distinct faults in Figure 3, with fault coordinates identifying the
 384 corresponding grid faces that act as fault boundaries. The upscaled fault leakage
 385 function as defined in Section 3.4 is integrated into these cells to account for fluid flow
 386 vertically along the fault. We assume a saline aquifer setting, where the initial reservoir
 387 pressure is set as hydrostatic, and the reservoir is fully saturated with brine. All
 388 simulations use a single injection well with an injection rate of 1 Mt per year for 30 years
 389 followed by a 970-year post-injection period, with open boundaries maintained at
 390 hydrostatic conditions. Total fault leakage is reported as the cumulative CO_2 flux that has
 391 broken through to the overlying aquifer at the fault top.

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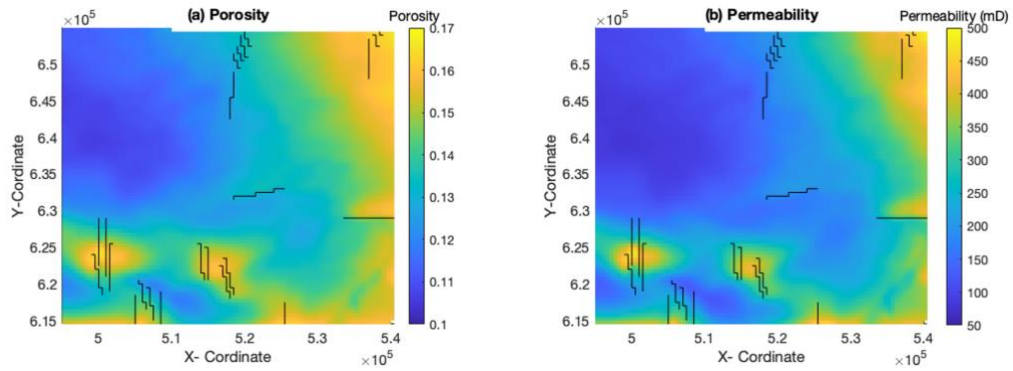


Figure 4 - Spatial distribution of reservoir properties calculated from depth. Black lines show the 24 mapped faults. (a) Porosity distribution ranging from 0.05 to 0.18. (b) Permeability distribution ranging from 100 to 550 mD. Both properties were calculated as functions of depth using petrophysical relationships shown in Eq. 8-9. The correlation between porosity and permeability distributions assumes depth-controlled diagenetic processes influencing reservoir characteristics (de Jonge-Anderson et al., 2024b).

393 Table 1 – Model parameters for Group J reservoir simulation in the Malay Basin. Porosity and permeability
 394 relationships (Eqs. 8-9) are from de Jonge-Anderson et al. (2025a). Fault parameters shown here represent
 395 the baseline Case A (sealed faults); variable fault properties for Cases A and B are given in Table 2.

Property	Values
Cell Dimensions (DX*DY) (m)	500 x 500
Cell thickness (DZ) (m)	100
Average top reservoir depth (m) (z)	2300
Rock compressibility (Pa ⁻¹)	4.35 x 10 ⁻¹⁰
Seafloor temperature (°C)	24
Temperature gradient (°Ckm ⁻¹)	50
Seafloor Depth (m)	70
Brine viscosity (Pa.s)	3.13 x 10 ⁻⁴
CO ₂ viscosity (Pa.s)	3.21 x 10 ⁻⁵
Brine density (kgm ⁻³)	1001.0
CO ₂ density (kgm ⁻³)	420.0
Fault (damage zone) permeability (mD)	See Table 2
Fault (damage zone) width (m)	5
Caprock thickness (m)	500
Fault (damage zone) Capillary Entry Pressure (bar)	See Table 2
Residual gas saturation (Reservoir)	0.2
Irreducible brine saturation (Reservoir)	0.3

396
 397 Table 2 – Fault property scenarios for well location analysis. Case A (sealed) represents structural
 398 containment with no fault leakage. Case B (leaking) represents compromised fault sealing with both
 399 vertical leakage along faults and reduced lateral sealing.

Case	Fault Capillary Entry Pressure (p_e) (bar)	Fault Permeability (k_f) (mD)	Fault Transmissibility Multiplier
Case A	∞	0	0
Case B	0.05	10 ⁻³	0.01

400
 401 4.1.2 Scenario Definitions

402 We analyse two scenarios to assess the impact of fault sealing capacity on storage
 403 performance (Table 2). Case A represents fully sealing faults where both horizontal
 404 (across-fault) and vertical (along-fault) flow are completely blocked, establishing the
 405 upper bound for CO₂ retention. Case B represents leaking faults with properties reflecting
 406 compromised sealing capacity: fault permeability of 10⁻³ mD (vertical flow along damage

407 zone), capillary entry pressure of 0.05 bar (reduced resistance to CO₂ entry), and
408 transmissibility multiplier of 0.01 (across-fault flow reduced to 1% of reservoir
409 permeability). In natural fault systems, damage zone width, permeability, and capillary
410 entry pressure are systematically related to fault displacement (Caine et al., 1996; Childs
411 et al., 2009; Faulkner et al., 2010). For this illustrative comparison, we assume uniform
412 fault properties across all 23 mapped faults. While this assumption is geologically
413 simplistic, it enables clear demonstration of how fault sealing capacity influences
414 migration patterns and storage security. Section 4.2 relaxes this assumption by treating
415 fault properties as uncertain parameters with spatially variable distributions.

416 4.1.3 Results

417 Figure 5 presents the baseline scenario where all faults are fully sealing in both horizontal
418 and vertical direction. CO₂ retention is defined as the total injected mass minus
419 cumulative boundary losses and fault leakage after 1000 years. Under Case A, lateral
420 boundary loss is the sole mechanism of CO₂ loss from the system. Retention varies
421 spatially across the 600 well locations (Figure 5a), controlled primarily by proximity to
422 domain boundaries and structural position. Wells placed within fault-bounded
423 compartments or in downdip structural depressions achieve higher retention than
424 boundary-proximal locations. Maximum overpressure (pressure above the initial
425 hydrostatic pressure) remains below 4 MPa across all well locations (Figure 5b),
426 indicating that injection at 1 Mtpa does not approach pressure limits that would constrain
427 operations at this site. The northeastern region, the deepest and least permeable part of
428 the domain (~50 mD and 10% porosity), experiences the highest overpressure despite
429 low retention. This is because, in low-transmissivity settings, reduced lateral pressure
430 dissipation results in localised pressure build-up.

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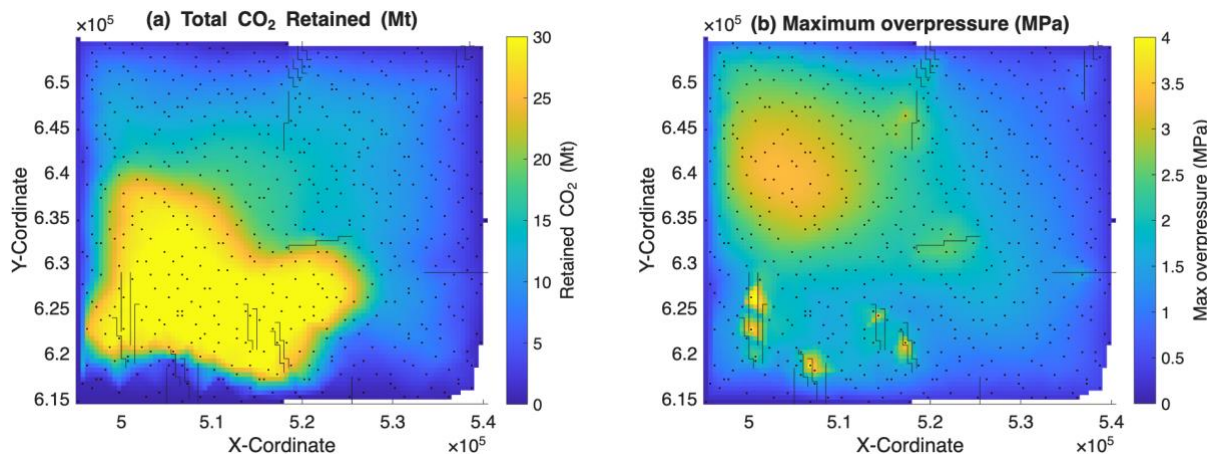


Figure 5 – Spatial distribution of CO₂ retention and pressure buildup under Case A (sealed faults baseline). (a) Total CO₂ retained after 1000 years for 600 simulated well locations (red dots), computed as total injected mass minus boundary losses and fault leakage. Faults (black lines) are fully sealing both laterally and vertically. High retention zones (warm colours) indicate locations with optimal structural containment, while low retention zones (cool colours) represent areas with greater lateral boundary losses. (b) Maximum overpressure above initial hydrostatic pressure observed anywhere in the reservoir over the 1000-year simulation period for the specific well location. All locations remain below 4 MPa, indicating pressure buildup stays within safe operational limits at 1 Mtpa injection rate.

432

433 Figure 6 examines the same well locations under compromised fault sealing conditions
434 (Case B). CO₂ retention (Figure 6a) shows spatially variable reduction relative to Case A,
435 with losses concentrated near faults intercepting the plume migration pathways. The

436 southwestern structural depression, bounded by major fault zones, limits lateral
 437 connectivity to domain boundaries, maintaining higher retention than locations with
 438 direct updip pathways to domain boundaries. Maximum overpressure (Figure 6b) exhibits
 439 similar spatial distribution to Case A, with peak values remaining below 4 MPa. CO₂
 440 leakage along faults (Figure 6c) reaches up to 0.3 Mt (~ 1% of injected mass) in locations
 441 where plumes directly contact faults, concentrated in the central and northern portions
 442 of the domain where structural dip drives updip migration toward faults. Brine leakage
 443 along faults (Figure 6d) precedes CO₂ plume arrival, driven by injection-induced
 444 overpressure at the fault base. Cumulative brine leakage reaches up to 2 Kt,
 445 approximately 150 times below the maximum CO₂ leakage of 0.3 Mt. This early brine
 446 leakage is nonetheless physically meaningful: the associated pressure dissipation at the
 447 fault reduces the injection-induced overpressure driving force (Eq. 6) when CO₂
 448 subsequently contacts the fault.

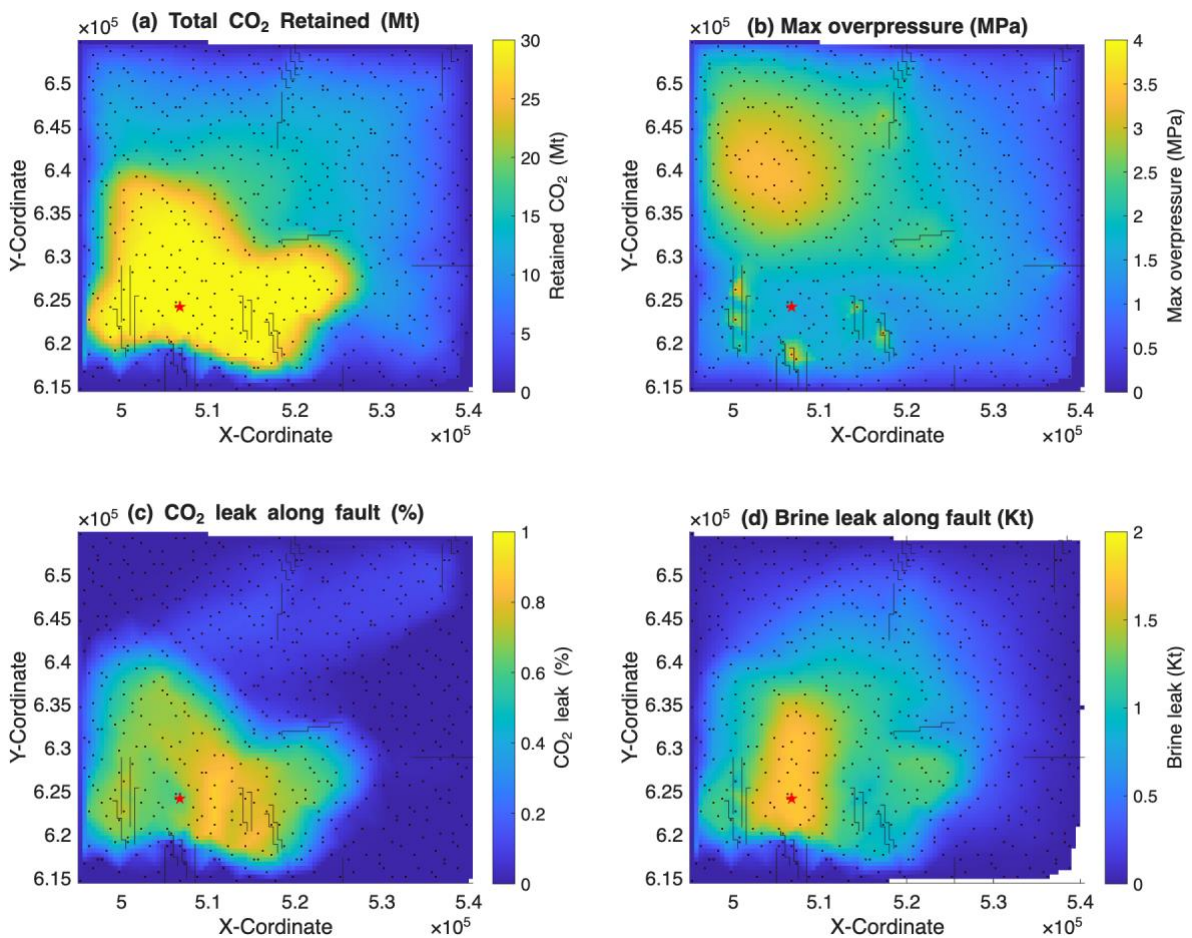


Figure 6 – Spatial assessment of CO₂ retention and fault leakage under Case B (leaking faults scenario). (a) Total CO₂ retained after 1000 years shows reduced retention compared to Case A. (b) Maximum overpressure distribution remains similar to Case A. (c) Cumulative CO₂ leakage along faults reaches up to 1% (0.3 Mt), concentrated where plumes migrate updip and contact fault planes. The cumulative leak is divided by the total injected to arrive at the percentage value shown here. (d) Cumulative brine leakage along faults reaches up to 2 Kt, driven by early-time injection-induced overpressure before CO₂ plume arrival. Red star marks the optimal location [506750, 624150], selected by sequential filtering on overpressure, boundary connectivity, and CO₂ retention. Black lines show faults; red dots show 600 well locations.

451 The optimal well location is identified using a sequential filtering approach. First,
 452 locations where maximum overpressure exceeds 2 MPa are excluded; this threshold is
 453 set as a conservative screening-stage proxy for geomechanical risk, set at half the
 454 maximum overpressure observed (Figure 6b); site-specific stress characterisation is
 455 required to define a geomechanically rigorous pressure limit. Second, locations where
 456 CO₂ migrates across the open model domain boundaries are excluded, as lateral
 457 boundary loss implies that containment depends on structural or hydraulic features
 458 outside the modelled domain. The remaining locations are ranked by CO₂ retention
 459 fraction (total injected mass minus boundary losses and fault leakage, normalised by
 460 total injected mass) after 1000 years, and the top-ranked location at coordinates
 461 [506750, 624150] is selected (red star, Figure 6). This location sits within the
 462 southwestern structural depression between two major fault zones, where structural
 463 compartmentalisation limits lateral connectivity to domain boundaries. Cross-
 464 referencing with Figure 4 reveals this region exhibits slightly higher porosity (~12-15%)
 465 and permeability (~300-400 mD). Storage patterns show a systematic decrease towards
 466 boundaries due to lateral migration across model boundaries. This structural depression
 467 provides favourable storage conditions despite being shallower than the deepest
 468 portions located in the northeastern part of the model domain. Locations near multiple
 469 fault intersections or in shallow updip positions show increases in leakage risk under
 470 Case B conditions (Figure 6c).

471 4.2 Uncertainty Analysis

472 4.2.1 Spearman Sensitivity Analysis

473 Unlike the deterministic Cases A and B that assumed uniform fault properties across all
 474 23 faults, real fault systems exhibit heterogeneous sealing capacity controlled by
 475 spatially variable permeability, capillary entry pressure, and transmissibility. To quantify
 476 the influence of fault-specific heterogeneity on leakage risk, we conducted 10,000 Monte
 477 Carlo simulations using Latin hypercube sampling across 71 uncertain parameters: 23
 478 fault permeability values, 23 fault capillary entry pressure values, 23 fault transmissibility
 479 multipliers, plus reservoir porosity and permeability exponents (Table 3). Fault properties
 480 were sampled independently for each fault from distributions spanning the ranges shown
 481 in Table 3. This permeability range (10^{-6} to 10^{-2} mD, equivalent to 10^{-21} to 10^{-17} m²) spans
 482 effectively unfractured caprock conditions at the lower bound to well-connected
 483 fractured damage zones at the upper bound, consistent with effective damage zone
 484 permeabilities derived from single-fracture stress-permeability relationships and
 485 fracture density distributions at natural CO₂ leakage sites (Snippe et al., 2022). Each
 486 simulation tracked CO₂ remaining in the reservoir, lateral boundary losses, and
 487 cumulative fault leakage over 1000 years. For each of the 71 parameters, we calculated
 488 the Spearman correlation coefficient (r_s) between the parameter value and total fault
 489 leakage across all 10000 realisations. Figure 7 presents the spatial distribution of fault
 490 sensitivity alongside a composite ranking of all parameters.

491
 492 Table 3 - Parameter distributions for uncertainty analysis. Fault parameters sampled independently for
 493 each of 23 faults. All fault properties follow log-uniform distributions; reservoir exponents (porosity and
 494 permeability) follow truncated normal distributions. All other parameters match Table 1.

Property	Distribution	Low Values	High Values
Fault permeability (mD) (x23)	Log-Uniform	10^{-6}	10^{-2}
Fault capillary entry pressure (bar) (x23)	Log-Uniform	0.01	10
Fault transmissibility multiplier (-) (x23)	Log-Uniform	10^{-6}	10^{-1}

Reservoir porosity exponent (-)	Truncated-Normal	0.000383	0.000684
Reservoir permeability exponent (-)	Truncated-Normal	3.4749	4.6264

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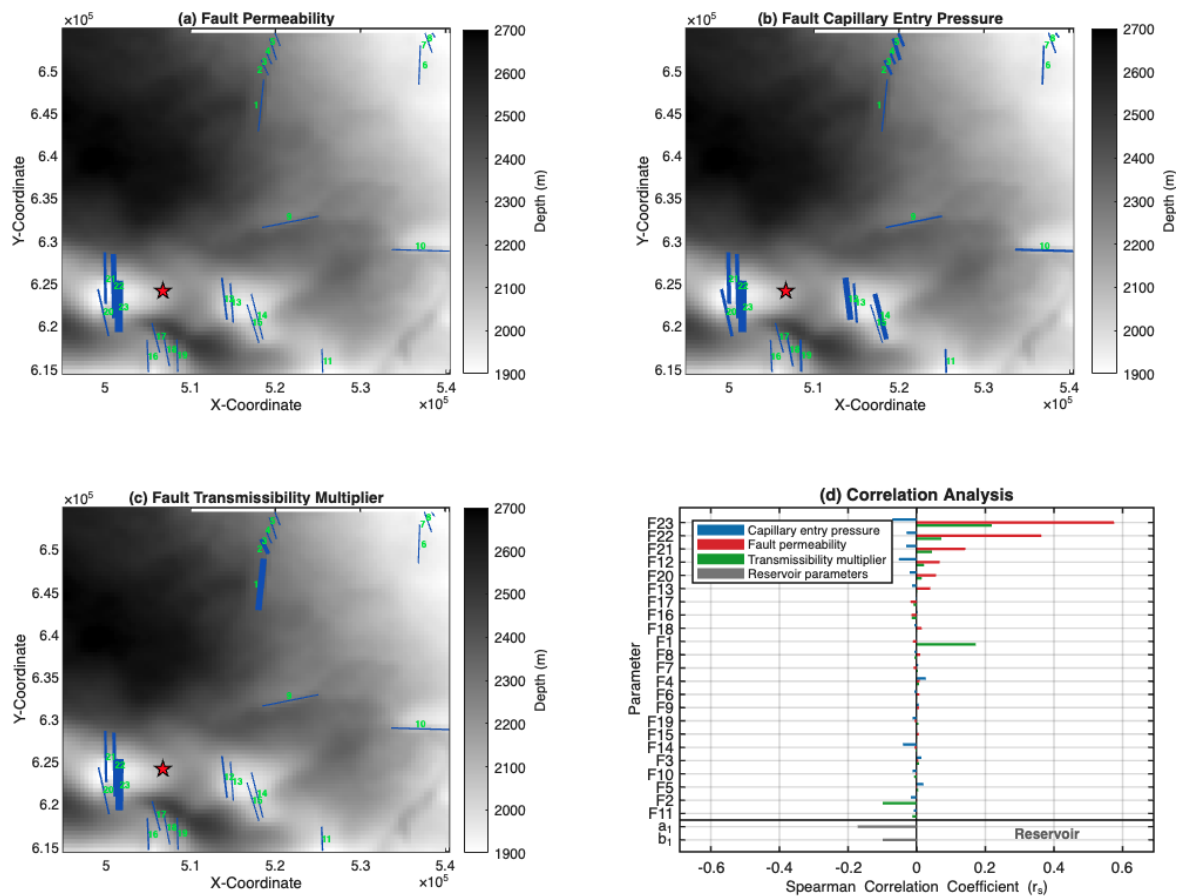


Figure 7 - Fault-specific sensitivity analysis from 10,000 Monte Carlo realisations. (a-c) show fault locations coloured and weighted by Spearman correlation coefficient with total fault leakage for permeability, capillary entry pressure, and transmissibility multiplier respectively. The fault is represented here in blue where the line thickness scales with correlation coefficient. Red star marks injection well location. Greyscale background is used to maximise contrast with fault line thickness, which scales with Spearman correlation coefficient. (d) presents composite correlation ranking for all 71 parameters, with faults sorted by correlation coefficient for fault permeability and reservoir parameters shown below.

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Leakage sensitivity concentrates on a small fault subset. Only 5 of 23 faults (Faults 23, 22, 21, 12, and 20) exhibit Spearman coefficients for the fault permeability > 0.05 against total leakage (Figure 7). These five faults represent 22% of the fault population yet dominate system uncertainty, providing a clear target where further site characterisation and data acquisition are needed to constrain uncertainties. Within this updip subset, Fault 23 exhibits the strongest sensitivity, approximately 5-10 times larger than the next most influential faults. The spatial pattern of influential faults reflects CO₂ migration pathways controlled by structural geometry and injection location. Faults 20-23 lie in the southwestern updip region close to the injection well (Figure 7a), positioned along the plume migration pathway where CO₂ accumulates at structural highs. These faults intercept migrating CO₂ and exhibit strong positive correlation with permeability and transmissibility, confirming that their hydraulic conductivity controls leakage rates once CO₂ contact occurs. In contrast, faults distant from the well show negligible influence on total leakage despite spanning the same parameter ranges. Reservoir properties exhibit

511 negligible correlation with fault leakage, approximately 3-6 times weaker than the
 512 dominant fault sensitivities (Figure 7d).
 513 Faults further away from the injection well show relatively stronger negative correlation
 514 with capillary entry pressure than nearby faults despite contributing less to total leakage.
 515 For example, Fault 12 (distant, northern position) exhibits $r_s(p_e) = -0.053$ compared to
 516 Fault 23's $r_s(p_e) = -0.071$, despite Fault 23's permeability correlation being 9 times
 517 stronger. This distance-dependent sensitivity reflects the diminishing pressure drive with
 518 distance from the injection well: in farther locations where overpressure is small,
 519 capillary entry pressure becomes the primary control on whether leakage takes place,
 520 whereas near-well faults are already breached in most realisations, as elevated injection-
 521 induced overpressure exceeds the capillary entry pressure regardless of its sampled
 522 value.
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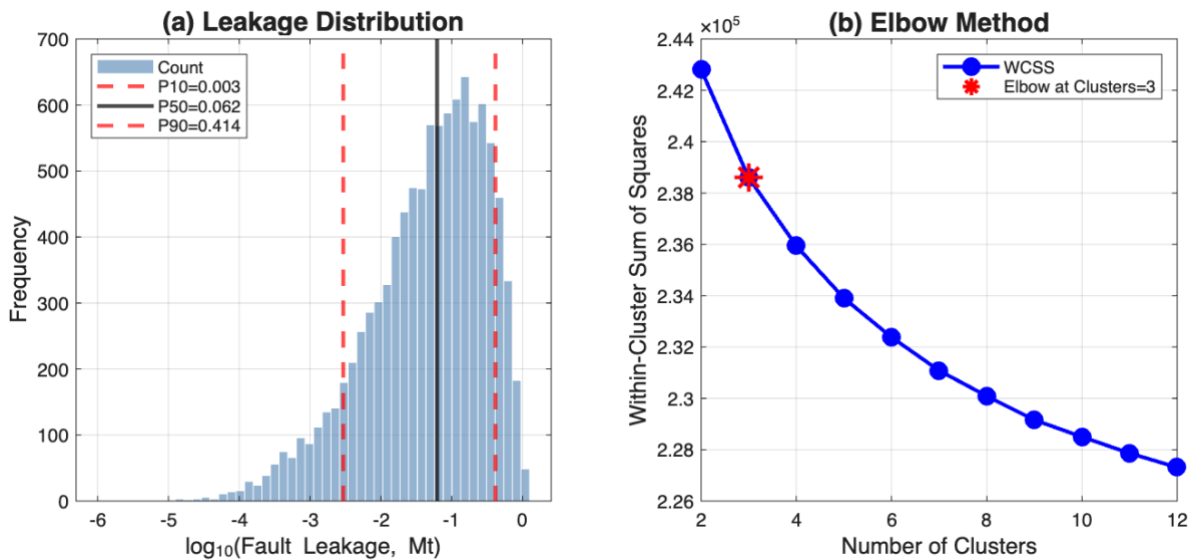


Figure 8 – (a) Probability distribution of total fault leakage after 1000 years from 10,000 Monte Carlo realisations. Vertical lines mark P10 (0.003 Mt), P50 (0.062 Mt), and P90 (0.414 Mt) quantiles. (b) Elbow method showing within-cluster sum of squares versus number of clusters, with optimal selection at 3 marked with a red star.

524 Unlike the other faults with significant Spearman correlation, which are controlled by
 525 fault permeability and capillary entry pressure as CO₂ migrates updip and leaks along the
 526 fault zone, Fault 1 is controlled by its transmissibility multiplier. Although Fault 1
 527 permeability shows minimal correlation with total fault leakage, its transmissibility
 528 multiplier exhibits substantial positive correlation with total fault leakage, ranking
 529 second among all faults for this parameter. This transmissibility dominance indicates
 530 Fault 1 functions as a "gateway fault": when its across-fault transmissibility is high, CO₂
 531 can migrate laterally through the fault zone and access the northern fault array (Faults 2-
 532 5), creating cascading leakage pathways. When Fault 1 maintains low transmissibility, it
 533 isolates the northern domain, preventing CO₂ from reaching downstream faults
 534 regardless of their individual properties. This finding cannot be identified through simple
 535 single-fault sensitivity analysis and demonstrates the value of comprehensive multi-fault
 536 uncertainty quantification.

537 4.2.2 Leakage Clustering Analysis

538 Total fault leakage after 1000 years spans approximately two orders of magnitude from
 539 P10 = 0.003 Mt (0.01% of injected mass) to P90 = 0.414 Mt (1.4% of injected mass) (Figure

540 8a). The median leakage of 0.062 Mt represents 0.21% of total injected CO₂. However, the
 541 upper tail of the distribution reaches P90 = 0.414 Mt, equivalent to 1.4% of injected mass,
 542 demonstrating that unfavourable combinations of fault properties can produce elevated
 543 leakage. The distribution exhibits strong positive skewness on a linear scale, with
 544 approximately 35% of realisations producing near-zero leakage (< 0.01 Mt) while less than
 545 10% exceed 0.2 Mt. This asymmetry indicates that high-leakage outcomes require
 546 specific combinations of fault properties rather than random parameter variations,
 547 motivating clustering to identify the parameter combinations governing distinct leakage
 548 regimes.
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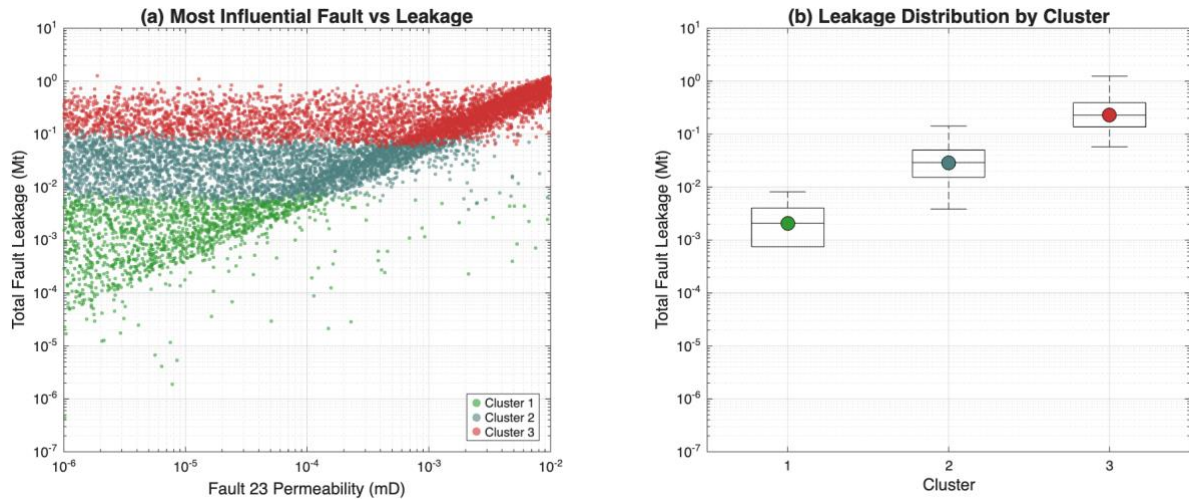


Figure 9 - Cluster optimization and characterisation. (a) Scatter plot of total fault leakage versus Fault 23 permeability (most influential fault) coloured by cluster assignment, demonstrating clear regime separation. (b) Box plots showing leakage distribution by cluster, with median values marked.

550
 551 K-medoids clustering partitions the 10,000 realisations into distinct behavioural regimes
 552 based on their positions in the 71-dimensional parameter space. The optimal number of
 553 clusters was determined using the elbow method (Figure 8b). The elbow occurs at cluster
 554 = 3, yields physically interpretable regimes (low, moderate, high leakage).
 555 Cluster 1 (low-leakage regime, n=1652, 16.5%) shows median leakage of 0.002 Mt
 556 (P10=0.0003, P90=0.006 Mt). Cluster 2 (moderate-leakage regime, n=3940, 39.4%)
 557 shows median leakage of 0.03 Mt (P10=0.010, P90=0.070 Mt). Cluster 3 (high-leakage
 558 regime, n=4408, 44.1%) shows median leakage of 0.230 Mt (P10=0.099, P90=0.588 Mt).
 559 The progression between clusters represents an ~15-fold increase from Cluster 1 to
 560 Cluster 2 and a ~8-fold increase from Cluster 2 to Cluster 3. The clear separation between
 561 regimes in leakage space reflects both the structure of the parameter distributions and
 562 the outcome-oriented weighting applied in the distance calculation (Section 3.5); the
 563 qualitative three-regime structure is expected to be robust to moderate changes in
 564 weighting, but the precise boundaries between clusters are conditioned on this choice.
 565 High-leakage scenarios (Cluster 3) represent the modal outcome, containing 44.1% of
 566 realisations compared to 16.5% in the low-leakage regime. Under current parameter
 567 uncertainty, elevated leakage is the single most probable outcome, exceeding the
 568 moderate-leakage regime by five percentage points. Cluster membership is
 569 predominantly controlled by a single fault (Fault 23) rather than distributed across the 23-
 570 fault system. Analysis of cluster-conditional parameter distributions for the top five
 571 influential faults (Faults 23, 22, 21, 12, 20) reveals that Fault 23 permeability exhibits a

572 ~100-fold difference between low and high-leakage clusters, substantially exceeding the
573 differences observed for other faults (35x for Fault 22, declining to 1.7x for Fault 20). The
574 visual separation of clusters in Fault 23 permeability space (Figure 9a) demonstrates that
575 this single parameter serves as an effective proxy for regime classification. The low-
576 leakage regime requires restrictive parameter combinations rather than simply avoiding
577 unfavourable values. Cluster 1 scenarios simultaneously maintain low permeability
578 across multiple updip faults (Faults 20-23) combined with elevated capillary entry
579 pressures, a narrow subset of the sampled parameter space.

580 Fault capillary entry pressure shows secondary influence on cluster assignment
581 compared to fault permeability. For the top five influential faults, the ratio of median fault
582 capillary entry pressure is 16–42% lower in the high-leakage cluster than in the low-
583 leakage cluster, indicating modest variation compared to the orders-of-magnitude
584 differences in permeability. This pattern is consistent with the correlation analysis
585 showing that capillary effects dominate in pressure-limited environments distant from
586 the injection well, whereas permeability controls leakage rates in the near-well region
587 where pressure drive remains elevated.

588 4.2.3 Conditional Cluster Probability

589 Figure 10 presents the conditional cluster probability for the top 15 parameters ranked by
590 Spearman correlation coefficient, showing the fraction of realisations in each leakage
591 regime as a function of parameter value. Dashed lines mark the unconditional cluster
592 fractions (C1: 16%, C2: 39%, C3: 44%); deviation from these baselines quantifies the
593 information each parameter value provides about leakage regime membership. Figure 10
594 identifies not only which parameters dominate leakage uncertainty, but what values they
595 need to take to shift cluster membership. Parameters whose curves deviate strongly from
596 the unconditional baselines are those for which a single measurement would
597 substantially resolve the current uncertainty; the crossing point where respective cluster
598 probability equals the unconditional fraction marks the threshold value separating
599 predominantly favourable from predominantly unfavourable outcomes.

600 Fault 23 permeability shows a response concentrated at the upper end of the sampled
601 range. Above approximately 10^{-3} mD, Cluster 3 conditional probability rises well above
602 the unconditional fraction of 0.44, approaching 0.9 near the upper bound. Below 10^{-3} mD,
603 this falls below 0.44 but does not approach zero, because Fault 22 permeability and Fault
604 1 transmissibility can independently drive high-leakage outcomes. The crossing point
605 near 10^{-3} mD therefore marks the value above which Fault 23 permeability alone indicates
606 elevated leakage risk, providing a concrete characterisation target that the Spearman
607 correlation coefficient-based ranking alone cannot supply. Fault 1 transmissibility shows
608 a different character: Cluster 3 conditional probability increases gradually from
609 approximately 0.35 to 0.60 across the full sampled range without a sharp crossing point,
610 indicating that no single measurement value decisively shifts cluster membership.
611 Instead, any reduction in transmissibility uncertainty progressively narrows the probable
612 leakage regime, consistent with its role as a gradual control on CO₂ access to the
613 northern fault array rather than an on/off sealing threshold. For the remaining parameters
614 in the ranking, conditional probability curves show progressively smaller deviations from
615 their unconditional baselines, confirming that below the top five (Fault 23 permeability,
616 Fault 22 permeability, Fault 23 transmissibility, Fault 1 transmissibility, and the porosity
617 exponent α_1), individual parameter values provide little additional information about
618 leakage regime membership beyond the prior.

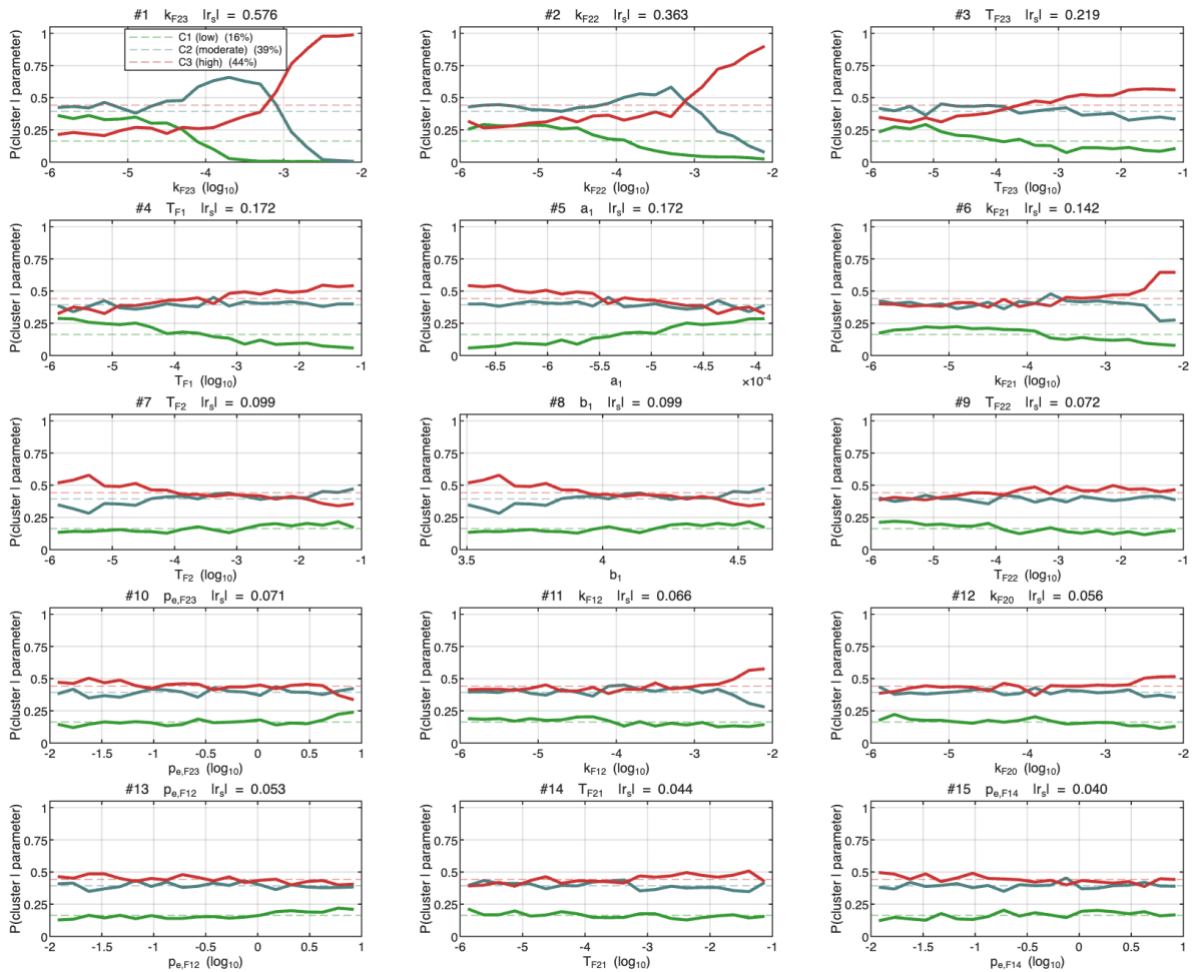


Figure 10 – Conditional cluster probability for the top 15 parameters ranked by Spearman coefficient. Each panel shows conditional probability as a function of the parameter, where the y-axis is probability (0–1) and the three solid lines (green: C1 low leakage, teal: C2 moderate leakage, red: C3 high leakage) sum to 1 at every x-value. Fault properties are plotted on a \log_{10} axis; reservoir exponents (a_1 , b_1) on a linear axis. Dashed lines indicate unconditional cluster fractions (C1: 16%, C2: 39%, C3: 44%), representing the probability prior to observing any parameter value. Solid-line deviation from the dashed baseline quantifies the discriminating power of each parameter: strong deviation indicates a high-value characterisation target; coincidence with the baseline indicates negligible influence on leakage regime.

620

621 The clustering identifies three medoid scenarios representing physically realisable
 622 parameter combinations for each regime. These medoid realisations can be extracted for
 623 detailed simulations, enabling assessment of effects neglected here without requiring
 624 full computational evaluation of the entire ensemble with 10,000 realisations. The
 625 conditional cluster probability analysis extends this further, identifying the specific
 626 parameter values at which cluster membership shifts and translating the correlation
 627 ranking into concrete characterisation targets. Together, the correlation, clustering, and
 628 conditional probability analyses shift risk assessment from parameter sensitivity
 629 analysis to scenario identification, enabling site selection that balances operational
 630 targets against geological constraints rather than applying blanket exclusion criteria.

631 5 Discussion

632 Structurally complex faulted basins have historically been regarded with caution as CO_2
 633 storage targets, yet the present results suggest that quantitative risk screening can

634 reframe this perception. Such basins were deprioritised due to perceived leakage risks.
635 Natural CO₂ accumulations show both outcomes at faulted sites: Bravo Dome and St
636 Johns Dome have retained large CO₂ volumes over geological timescales despite active
637 faulting, while fault-focused surface seepage has been documented at Green River, Utah,
638 where CO₂ migrates along normal fault damage zones (Miocic et al., 2019a, 2019b;
639 Snippe et al., 2022). The same framework is applicable to subseismic faults that lie below
640 seismic resolution; their potential presence and properties can be incorporated as
641 additional uncertain parameters within the Monte Carlo framework, enabling leakage risk
642 to be bounded where fault mapping is incomplete. The coexistence of these outcomes
643 across mapped and potentially unmapped fault populations suggests that faulted basins
644 are not uniformly unsuitable for storage; quantitative fault characterisation represents a
645 more appropriate early-stage response than blanket avoidance.

646 Well placement in faulted storage systems requires balancing structural confinement
647 against pressure management constraints. Pressure propagation extends far beyond the
648 CO₂ plume at basin scale (Birkholzer & Zhou, 2009; Zhou & Birkholzer, 2011), and
649 pressure space, the capacity of connected formations to absorb injection-induced
650 overpressure without triggering fault reactivation or injectivity loss, is increasingly
651 recognised as the primary limitation on storage scalability (Birkholzer & Zhou, 2009;
652 Bump & Hovorka, 2024; Ougier-Simonin et al., 2026). Well placement optimisation and
653 brine extraction strategies have been proposed to address this constraint (Cihan et al.,
654 2015; González-Nicolás et al., 2019), with recent work extending to fault slip minimisation
655 as an explicit objective (Raji et al., 2025) and to basin-scale pressure assessment in
656 multi-fault systems (Cihan et al., 2026). The sequential filtering approach applied here
657 addresses the same coupled problem at early screening stage: in this study,
658 overpressure and boundary connectivity together constrain the viable well location
659 space more effectively than optimising on CO₂ retention alone, with the southwestern
660 structural depression emerging as the preferred location. This structural depression
661 warrants further characterisation, as fault-bounded compartmentalisation limits lateral
662 connectivity while maintaining adequate transmissivity for pressure dissipation (Zhou et
663 al., 2008; Mears et al., 2026). This result is consistent with observations that structural
664 position and boundary conditions collectively control effective storage volumes in
665 compartmentalised systems (Mhlambi et al., 2025), and that low-transmissivity settings
666 can simultaneously constrain injectivity and amplify local pressure buildup (de Jonge-
667 Anderson et al., 2026).

668 The dominant influence of fault permeability and capillary entry pressure on leakage
669 outcomes carries a direct implication for characterisation programme design that
670 extends beyond the simple observation that both parameters are uncertain. Capillary
671 entry pressure governs the onset threshold for vertical CO₂ flow: faults with sufficiently
672 elevated entry pressure remain effectively sealing even under injection-induced
673 overpressure, because the combined buoyancy and pressure driving potential cannot
674 overcome the capillary resistance at the fault base (Espinoza & Santamarina, 2017;
675 Busch & Kampman, 2018). The relative importance of entry pressure and permeability is
676 not, however, spatially uniform. Faults in close proximity to the injection well, in our
677 simulations, experience injection-induced overpressure sufficient to breach capillary
678 resistance across a wide range of entry pressure values, placing permeability as the
679 dominant control on leakage magnitude in this regime. Faults distant from the injection
680 well, in our simulations, experience lower injection-induced overpressure due to

681 pressure dissipation with distance; at these locations, capillary entry pressure becomes
682 the primary determinant of whether leakage initiates at all. This spatial transition is
683 directly evidenced in the correlation analysis: Fault 23 exhibits permeability sensitivity
684 approximately an order of magnitude larger than its capillary entry pressure sensitivity,
685 while farther faults such as Fault 14 and Faults 2-5 show relatively stronger capillary
686 sensitivity despite weaker aggregate influence on total leakage. This spatial distinction
687 also has a temporal dimension: faults in close proximity to the injection well are exposed
688 to peak injection-induced overpressure during the 30-year injection period and are
689 therefore likely to leak predominantly during operations, whilst faults distant from the
690 well are reached only after CO₂ has migrated updip during the post-injection period,
691 representing a longer-term post-injection risk horizon. A characterisation programme
692 can be sequenced accordingly, prioritising permeability measurement for the updip fault
693 subset where overpressure drive is sustained, while capillary entry pressure
694 measurements constrain leakage onset thresholds for all faults. In practice, both would
695 rely on analogue fault data from similar clastic settings where direct core measurements
696 are unavailable (Faulkner et al., 2010; Bense et al., 2013; Busch & Amann-Hildenbrand,
697 2013; Espinoza & Santamarina, 2017).

698 The concentration of leakage along faults intercepting the updip migration pathway is
699 consistent with field-scale studies showing that leakage outcomes concentrate in a
700 small structural subset rather than distributing uniformly across fault populations
701 (Gasda et al., 2022; Salo-Salgado et al., 2025). Only 5 of 23 faults exhibit significant
702 correlation with total leakage, all positioned along the primary plume migration pathway,
703 which suggests that structural geometry provides a practical first-order guide for focusing
704 characterisation investment even before detailed fault property data are available. Fault
705 23 dominates system response consistent with its position at the primary updip migration
706 pathway, and targeted characterisation of this fault alone would substantially reduce
707 current predictive uncertainty relative to uniform sampling across the full fault
708 population. The gateway behaviour of Fault 1 is a qualitatively different result that
709 structural position cannot predict: its transmissibility controls lateral access to an entire
710 northern fault array, such that low transmissibility isolates this domain regardless of the
711 individual properties of faults downstream. This cascade pathway mechanism is invisible
712 to single-fault sensitivity analysis, which would assign negligible influence on Fault 1
713 based on its permeability correlation alone, and demonstrates that multi-fault
714 uncertainty quantification is necessary to resolve system-level leakage behaviour in
715 structurally complex settings. Recent approaches have advanced fault leakage
716 uncertainty quantification, applying full two-phase physics and geomechanical coupling
717 to single-fault systems (Lu et al., 2025) or copula representation of dependent fault
718 properties within a single fault zone (Pettersson et al., 2025); the present study
719 complements these by resolving the cascade interactions that emerge across a 23-fault
720 population simultaneously. The conditional cluster probability analysis translates this
721 ranking into concrete characterisation targets, identifying the specific parameter values
722 at which leakage regime membership shifts and providing actionable measurement
723 thresholds that the Spearman ranking alone cannot supply.

724 The 44% modal frequency of high-leakage realisations reflects the breadth of current
725 parameter uncertainty rather than an inherent property of the storage site. As fault
726 properties are constrained through targeted characterisation, parameter uncertainty
727 narrows and realisations shift toward the cluster consistent with measured values; this

728 may confirm lower leakage risk or reveal that high-leakage conditions are more probable
729 than current parameter distributions suggest (Alcalde et al., 2018). The three medoid
730 scenarios are specific, physically realisable parameter combinations that can be used
731 directly as inputs to detailed 3D coupled simulations, enabling assessment of effects
732 excluded here without evaluating the full ensemble (Scheidt & Caers, 2009; Mahjour &
733 Faroughi, 2023).

734 CO₂ movement to the overlying aquifer via fault damage zones is classified as leakage in
735 this analysis, as it exits the defined storage formation into a separate compartment.
736 Under typical regulatory frameworks, such flux would instead be classified as migration
737 if the overlying aquifer were formally designated within the storage complex boundary
738 (Dixon et al., 2015; Krevor et al., 2023). The distinction carries regulatory weight: leakage
739 implies loss of containment assurance, whereas within-complex migration implies
740 redistribution within a characterised volume. For early-stage screening, the conservative
741 interpretation is appropriate; storage complex boundary definitions would be established
742 as part of the permitting process for any project advancing toward the permitting stage.
743 The analysis treats fault properties as uniform within each fault but heterogeneous
744 across the 23-fault population, and as temporally constant throughout the simulation
745 period. Real fault zones exhibit spatial variability along strike and depth, with preferential
746 leakage pathways developing through clay-smear discontinuities and damage-zone
747 variability (Berge et al., 2022; Snippe et al., 2022; Salo-Salgado et al., 2023; Rizzo et al.,
748 2024). Within-fault heterogeneity was excluded to maintain tractability across 10,000
749 realisations; the independent sampling of properties across 23 faults partially addresses
750 between-fault variability but does not capture co-variation between structurally related
751 faults that copula-based approaches can resolve (Pettersson et al., 2025). The model
752 neglects geomechanical coupling and non-isothermal transport effects, assuming
753 instead that CO₂ has equilibrated thermally with formation temperature prior to reaching
754 the fault zone. Geomechanical effects could modify leakage through stress-induced
755 permeability changes and fault reactivation (Rutqvist, 2012; Ellsworth, 2013; Rinaldi et
756 al., 2014; Vilarrasa et al., 2019; Zhao and Jha, 2019; Chen et al., 2024; Zheng et al., 2025).
757 Recent probabilistic fault stability assessment emphasise the importance of fault throw
758 relative to reservoir thickness and pore pressure change in evaluating reactivation risks
759 (Yoon et al., 2023; Bisdorn & Chan, 2024), and integrated mechanical and dynamic fault
760 analyses would complement the present screening approach for sites where
761 geomechanical risk is elevated. Reservoir property uncertainty influences CO₂ migration
762 rates and plume geometry, and therefore which faults are contacted and when; this is
763 captured partially through the reservoir exponent distributions but not through spatially
764 variable porosity and permeability fields. The k-medoids clustering includes total fault
765 leakage as a weighted variable in the distance metric, orienting cluster boundaries
766 toward leakage behaviour rather than pure parameter-space proximity; while the
767 qualitative three-regime separation is expected to be robust to moderate changes in this
768 choice, the precise cluster boundaries and medoid positions are conditioned on the
769 weighting applied, and systematic testing of alternative schemes remains a direction for
770 methodological refinement.

771 The computational efficiency of the approach, with each simulation completing in under
772 30 seconds, is what makes the 10,000 realisations tractable as an early-stage screening
773 tool rather than a dedicated research exercise. This efficiency is particularly
774 consequential in settings such as the Malay Basin, where Malaysia's CCS deployment

775 targets require systematic evaluation of fault leakage risk across multiple aquifer
776 intervals and structural configurations within compressed timescales (Fernandez et al.,
777 2024; de Jonge-Anderson et al., 2025a). The workflow demonstrated here, combining
778 rapid fault leakage screening with ensemble uncertainty quantification, Spearman
779 sensitivity analysis, k-medoids clustering, and conditional cluster probability analysis,
780 provides a transferable framework for early-stage risk characterisation in any faulted
781 storage setting where comprehensive fault property data are unavailable at the time of
782 site screening.

783 6 Conclusions

784 This study demonstrates rapid risk assessment for CO₂ storage in structurally complex
785 basins. Application to the 23-fault system in the Malay Basin's Group J aquifer shows that
786 computationally efficient screening tools enable systematic evaluation of well locations
787 and fault property uncertainties. Evaluation of 600 well locations using sequential
788 filtering on overpressure, boundary connectivity, and CO₂ retention identified an optimal
789 injection site in the southwestern structural depression, where fault-bounded
790 compartmentalisation limits lateral boundary losses while maintaining adequate
791 transmissivity for pressure dissipation. Uncertainty quantification across 71 parameters
792 using 10,000 Monte Carlo realisations reveals three key findings. First, leakage sensitivity
793 concentrates spatially: only 5 of 23 faults exhibit significant correlation with outcomes,
794 indicating that targeted characterisation of the updip fault subset would substantially
795 reduce predictive uncertainty, rather than requiring comprehensive characterisation
796 across the full fault population. Second, high-leakage scenarios comprise 44% of
797 realisations under current parameter uncertainty, representing the modal outcome
798 rather than a tail event, and reflecting that the current parameter distributions are
799 insufficiently constrained to distinguish favourable from unfavourable outcomes. Third,
800 cluster membership is predominantly controlled by Fault 23 permeability, exhibiting
801 ~100-fold contrast between regimes, while Fault 1 transmissibility controls access to an
802 entire downstream fault array, a cascade pathway that structural position alone cannot
803 predict and single-fault analysis would miss. Conditional cluster probability analysis
804 translates this ranking into concrete characterisation targets, identifying the specific
805 parameter values at which leakage regime membership shifts, information the
806 correlation ranking alone cannot supply. Together, these findings shift risk assessment
807 from parameter sensitivity analysis to scenario identification, enabling site selection that
808 balances operational targets against geological constraints. The computational
809 efficiency, with each simulation completing in under 30 seconds, enables
810 comprehensive uncertainty quantification supporting targeted data acquisition during
811 early-stage screening.

812 Data Availability Statement

813 Simulations were conducted using co2-lab module of the open-source MATLAB Reservoir
814 Simulation Toolbox (MRST), available at <https://www.sintef.no/projectweb/mrst/>. The
815 fault leakage function module is on GitHub (<https://github.com/Poriyalar/faultve>) and
816 available at Zenodo (<https://doi.org/10.5281/zenodo.17469206>). The Group J grid and an
817 example script for loading it and mapping faults are available at Zenodo
818 (<https://doi.org/10.5281/zenodo.21137719>).

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824 AI statement

825 During the preparation of this work the author(s) used Claude in order to improve the
826 readability. After using this tool/service, the author(s) reviewed and edited the content as
827 needed and take(s) full responsibility for the content of the published article.

828 CRediT Statement

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