

Biological Dynamics for Sustainable Carbon Sequestration

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Abstract

Earth's atmosphere is shaped by the interplay of biological systems operating across interlinked biogeochemical cycles and, more recently, by anthropogenic emissions. Here, we thoroughly examine the potential for engineering biological systems across ecosystems for sustainable carbon storage. This centralization and comparative analysis of the potential of all key global ecosystems aims to facilitate the integration of nature-based carbon sequestration solutions into the portfolio of solutions currently being explored. Beyond reporting the potential for long-term carbon storage of the various ecosystems present globally, we show that biological equilibria strongly influence biological carbon dynamics. Consequently, routes for the strategic implementation of biological carbon sequestration in out-of-equilibrium systems are put forward. As carbon sequestration by ecosystems is typically linked to other benefits such as food production and expansion of their service value, we expect that the expansion and integration of by-design biological carbon sequestration could significantly contribute to the management of atmospheric carbon.

Introduction

The sustained rise in atmospheric carbon dioxide (CO₂) levels has galvanised the search for effective carbon management solutions^{1,2}. Technological approaches such as carbon capture and storage (CCS)³, where carbon (C) is captured at the source and stored, and renewable energy deployment have both received major investments (exceeding 11 billion USD in 2024⁴ and projected to reach 100 billion per year by 2035⁵). In addition to CCS, technologies for carbon capture and utilisation (CCU),

where carbon is reused, and CO₂ removal (CDR), where CO₂ is actively removed from the atmosphere, have received much attention⁶.

Key sectors remain hard to abate, with energy and land-use collectively contributing to >70% and 15% of anthropogenic emissions, respectively (**Figure S1**)⁷. Though CDR programs are rapidly growing (750% year-on-year increase in 2022-2023)⁸, the IPCC estimates that 100-1000 Gt of CO₂ will need to be removed from the atmosphere by 2100 to limit global warming by 1.5°C⁹. Removal goals are proposed by the IPCC through a diverse portfolio of methodologies spanning Earth systems (land, air and ocean)⁶ comprises the carbon management landscape (CCS, CCU and CDR, collectively). Despite significant efforts, however, the scalable success of engineered approaches to carbon management has been limited³. Conversely, the natural biological processes regulating the global biogeochemical carbon cycle (**Figure 1**) remain underutilised, even though their yearly carbon exchanges are far exceeding those associated with anthropogenic emissions^{8,12,13}. Harnessing biological dynamics, therefore, presents a major opportunity for carbon management. However, there is presently no centralised information on the role of biological systems in carbon sequestration strategies and the associated potential of each ecosystems for long term carbon storage. Here, we synthesise the current knowledge on biological systems in the carbon sequestration landscape and identify strategic pathways for leveraging natural systems for scalable carbon management. By promoting the integration of biological systems into strategic CDR development, we aim to advance the development of sustainable sequestration solutions that align with food production and other ecosystem services.

Overview of carbon capture and global dynamics

Anthropogenic carbon emissions presently exceed the rate at which natural systems integrate atmospheric carbon (**Figure 1**). Indeed, the atmospheric C stock (~910 gigatonnes C (Gt C)), experiences an annual increase of ~0.46% yr⁻¹ (**Figure 1**), or cumulative increase of 235.51 Gt C between 1958 and 2025 (per NOAA direct measures, available at <https://climate.nasa.gov/vital-signs/carbon-dioxide/>), highlighting the need to improve our integration with natural systems. Anthropogenic emissions from the geological C reservoir (4,130 Gt)¹⁴ are primarily attributed to the combustion of fossil fuels at a rate of ~7.0 Gt C yr⁻¹ as CO₂ release. Engineered CCS have progressed from the basic strategies such as absorption and catalytic or non-catalytic solvent regeneration¹⁵⁻¹⁷ from pollution point sources, to direct air capture (DAC) that extracts CO₂ directly from the atmosphere^{18,19} and sequestration into underground geological formations (estimated capacity of 5,000–10,000 Gt) by injecting CO₂²⁰. However, as of 2023, engineered CDR systems sequester around 0.0045 Gt C annually²¹, less than 0.01% of global carbon exchanges (**Figure 1**) and 0.4% of the target required to meet the Paris Agreement goals by 2050²². In comparison, at 57.3-66.5 Gt C yr⁻¹, net primary production from terrestrial photosynthetic processes account for the largest annual drawdowns of atmospheric carbon, equating to ~13,000× of that achieved with CDR programs²³. Leveraging, or at least incorporating, the dynamics of biological systems will therefore become increasingly important. For example, given the magnitude of biological dynamics on carbon cycling, shifting them by 0.1% would result in ~13× the cumulative reductions

achieved with engineered CDR approaches. Critically, these processes can be further leveraged via, and integrated into, carbon credit systems if their cost-effectiveness is warranted (Supplementary **Discussion S1**).

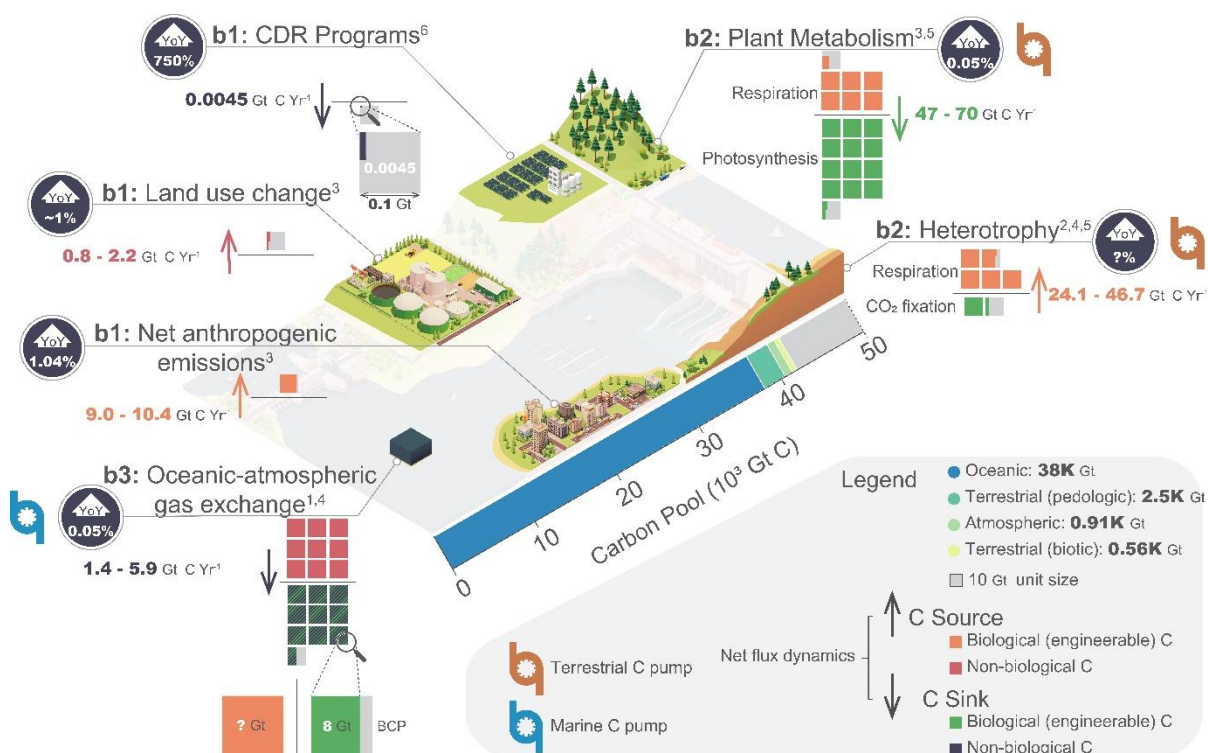


Figure 1 | Major landmarks of the global carbon (C) cycle. For each landmark (CDR Programs, Plant Metabolism, Land use change, Oceanic-atmospheric gas exchange, Heterotrophy and Net anthropogenic emissions). These landmarks are grouped by anthropogenic aspects (b1), terrestrial aspects (b2) and oceanic (b3). The latest year-on-year (YoY) change is presented (as the % against the prior year on record), and net flux is presented explicitly, as the latest ranges in terms of gigatonnes (Gt) per year. Adjacent arrows highlight whether the associated landmark is an overall C sink or source for atmospheric CO₂. Blocks accompanying arrows that are grouped above (CO₂ emission facets) and below (CO₂ absorption facets) the line represent gross units of 10 Gt C (except for CDR programs for which a 0.1 Gt unit is used for clarity) to facilitate comparison of the magnitude of main C landmark facets (for the BCP portion, this line has been made horizontal for illustrative purposes), for oceanic-atmospheric CO₂ absorption the green-blue dashed boxes represent our uncertainty as to the relative contributions of biological/non-biological processes. Blocks are coloured to highlight biological C cycling, which presents opportunities for long-term engineering of the biogeochemical C cycle to favour low atmospheric C. The principal carbon pools (atmospheric, oceanic, and terrestrial (split into pedologic & biotic fractions)) are presented as a stacked bar; explicit values for each pool are given in the legend. Where: GHG – greenhouse gases, BCP – Biological Carbon Pump and CDR – Carbon Dioxide Removal. The year for supporting data for each landmark is indicated where ¹ – 2012, ² – 2019, ³ – 2020, ⁴ – 2021, ⁵ – 2022 and ⁶ – 2023. Further information on data sources can be found in **Discussion S7**.

The four principal pools associated with the biogeochemical carbon cycle (**Figure 1**) are the *oceanic*, *terrestrial (pedologic)*, *atmospheric* and *terrestrial (biotic)* pools. The oceanic pool is the largest (with an estimated 38,000 Gt C), primarily (>90%) stored as dissolved inorganic carbonates (DIC)²⁴. Within this pool, the ocean-atmosphere gas exchange accounts for a significant portion of the biogeochemical carbon cycle, playing a critical role in regulating atmospheric CO₂ levels²⁵. This exchange is mediated by partial pressure differences between the gas and liquid phases and further shaped by the biological conversion of DIC into particulate inorganic carbon

(PIC). The pelagic transformation of DIC to PIC is intrinsically tied to calcifying marine organisms (notably coccolithophores and foraminifera), which contribute as much as 56% PIC production^{26,27}.

For terrestrial carbon, the pedologic (soil) pool carries the largest stock of carbon (~2,500 Gt C)²⁸, accounting for around three times the carbon stored in the atmosphere²⁹ (**Figure 1b**). Approximately 60% (~1,500 Gt C) of the pedological pool exists as soil organic carbon (SOC)³⁰ and is therefore derived from biological systems. The remaining 40%, the soil inorganic carbon (SIC), is also thought to be significantly influenced by biological processes, though the relative contributions of these processes across geographies, such as biomineralisation, bioturbation and mineral dissolution, is not yet clear^{28,31,32}.

The other primary terrestrial pool is the biotic pool (**Figure 1b**), the above-ground fraction of which is the largest and is primarily represented by plant life (comprising ~450 Gt C of a ~550 Gt C biomass pool size³³). The subsurface fraction (which is comprised of necromass, root systems, fungi, and microbial communities) is smaller than its topside counterpart by several orders of magnitude greater than the aboveground fraction stored in animal and microbial biomass. So, although animal biomass is critical for carbon cycling, it is less well suited for CCS. Notably, the dynamics and relative importance of the natural carbon pools are currently based on broad estimates. Improved quantification of these aspects would clarify the comparative roles of biological and engineered carbon sequestration strategies, supporting the optimisation of more targeted interventions and the development of effective nature-based solutions.

Carbon sequestration lifetime and its association with biological dynamics

Carbon sequestration strategies can be described in terms of six fundamental categories defining storage in 1) waste materials, 2) plant biomass, 3) soils, 4) man-made products, 5) oceans and 6) the deep subsurface (geological) (**Figure 2**). The corresponding BCS actors (including microbes, plants, fungi, and algae), play a fundamental role in many of these strategies (**Figure 2, right**). Their involvement brings both opportunities and challenges, influencing the effectiveness and longevity of CDR and long-term storage. Across these categories of carbon sequestration strategies, each of the products has its own recalcitrance to re-emission, which is proportional to their mean retention time (MRT) (**Figure 2, left**). Broadly speaking, waste biomass utilisation of durable goods and storage in certain man-made products (fuels, textiles, engineered woods, bioplastics and composites) possess the shortest MRTs (with reported MRTs often below a year)^{34–36}. Of note, there are currently efforts to increase the portion of materials formed from natural carbon sequestration - wood and plant-based materials, for example (**Discussion S2**).

These are followed by landfill sequestration, agricultural approaches, wood construction and ocean fertilisation (ocean fertilisation is notable for its high variability and context dependency and may therefore be underrepresented), where MRTs extended into the order of decades^{37,38}. Whilst they are therefore unsuitable for long-term sequestration initiatives, these approaches may yet find utility at local scales as part of holistic frameworks for carbon management. Long-term biomaterial-based (bioconcrete and biochar) methods, ocean alkalinity enhancement, and forestation

approaches begin to approach desirable sequestration timescales (~100s of years)^{39–41}. For truly long-term storage (millennia-Ma timespans), methods for C storage in sediments (such as blue carbon ecosystem (i.e., mangroves, salt marshes and seagrasses), and macroalgae cultivation with sinking) and low-turnover peatland soils, mineralisation or deep within geological formations are the primary avenues available^{39,42–44}. The geographical and temporal scales surrounding MRT calculations render accurate MRT assessment challenging, as direct observation is impractical. Current long-term MRTs are therefore necessarily limited by their dependence on modelled lifetimes. Nevertheless, at such time horizons, these latter approaches warrant special focus for development. However, it is also important to note that MRT is not the only consideration (an elaborated discussion of CCS MRTs is provided in **Discussion S3**, Supplementary Information), carbon costing, storage capacity, environmental, socioeconomic, and ethical considerations need to be factored into implementation planning. In that regard, each of the available approaches carries its own advantages and challenges.

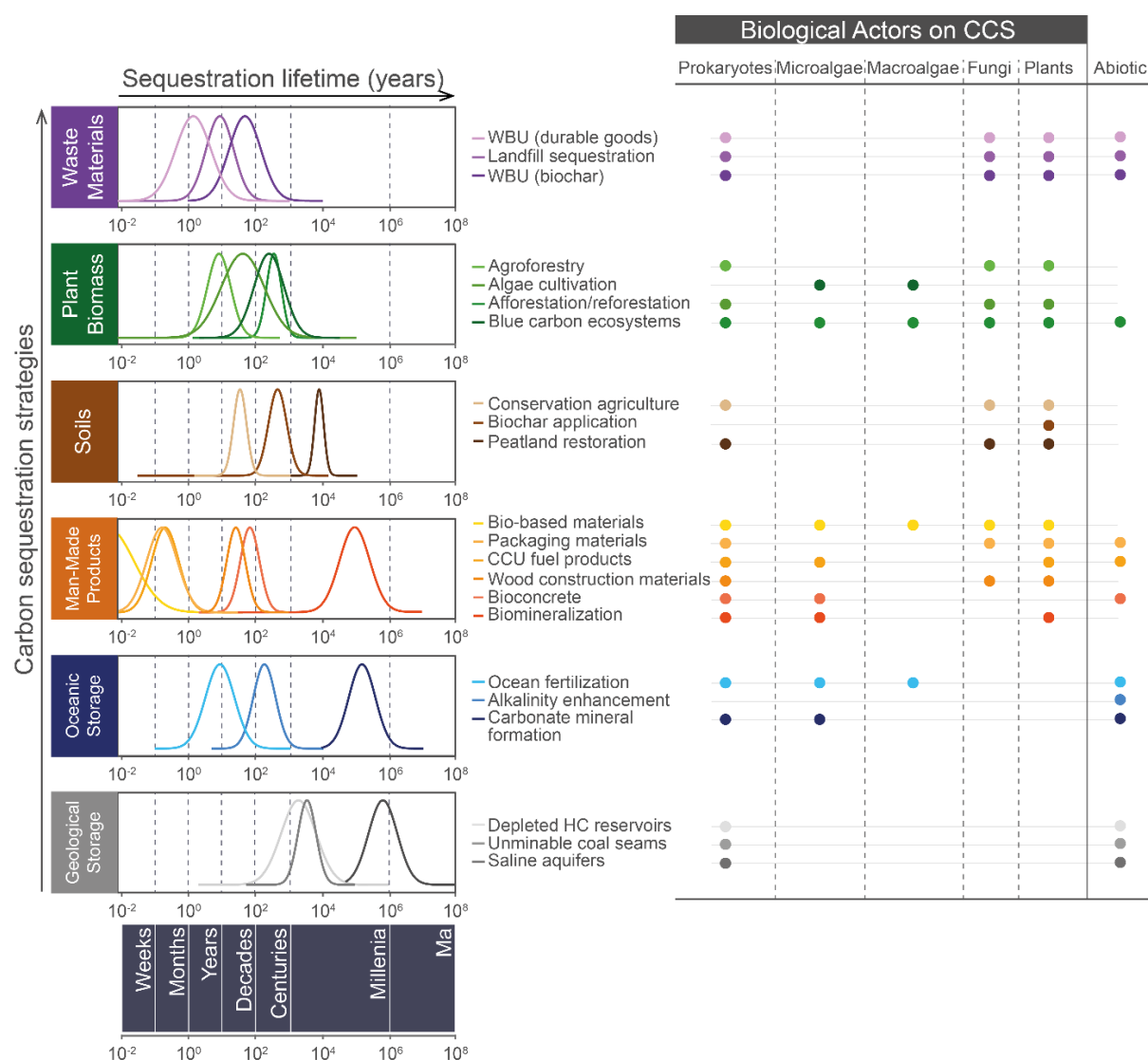


Figure 2 | Sequestration lifetime across ecosystems and anthropogenic sectors as associated to biological systems. Right panel - The representative normalised probability densities (log-normal distributions) for carbon retention lifetimes of carbon sequestration strategies (grouped by shared

approaches). The dotted lines indicate the time horizon, as shown in the bottom-most plot (indigo). Where WBU – waste biomass utilisation; HC - hydrocarbon; CCU carbon capture and utilisation. **Left panel** - Table indicating the involved biological systems (prokaryotes, microalgae, macroalgae, fungi or plant) that are either actively contributing to, or otherwise influencing, the associated sequestration strategy (as indicated by the coloured circles). Circles are also shown for strategies in which abiotic processes are the fundamental aspect. Further information on the data processing for this panel can be found in **Discussion S3**.

Accounting for the net carbon balance for natural systems remains especially challenging due to their complexity and scale. Life cycle assessments (LCA) have become an increasingly popular approach (influencing policy-making, industry, academia) to evaluate carbon sequestration⁴⁵. Until fairly recently, there was a consensus in carbon accounting and LCA modeling that carbon-neutral and climate-neutral were analogous terms. This paradigm accepted the idea that a net flux balance of zero emissions can be reached within a given system if the CO₂ absorbed equals the CO₂ emitted. However, by adopting a dynamic LCA approach⁴⁶, and applying it to a system where C rotates between the biosphere, technosphere and atmosphere (e.g., from biomass) the significance factoring in CO₂ molecules that remain in the atmosphere before being incorporated into biomass was revealed⁴⁷. For example, examining the atmospheric decay of a biogenic CO₂ pulse stemming from biomass with rotation and storage periods of 50 and 20 years, respectively (which is then released as a pulse), yields a net area under the curve (corresponding to its global warming potential) that should not be disregarded.

Therefore, it is essential to understand the dynamics of the system and relate how carbon storage in the technosphere, or delayed emissions contribute to climate change⁴⁷. In this sense, the “biogenic global warming potential” (GWP_{bio}) index was introduced to relate biomass growth, emissions from biodynamic fluxes, and the global carbon cycle^{47,48}. The characterization factor developed served to quantify the benefits or burdens of storing biomass in the technosphere and the subsequent C release with the radiative forcing integrated within a fixed time window. For instance, for the above example (rotation period of 50 years, 20 years of storage), the GWP_{bio} (for a 100-year time horizon GWP) equals to 0.05, meaning that 5% of the embedded carbon should be considered as emissions. This indicates that despite the storage, a portion of the embedded carbon can impact climate change. We believe this insight is valuable when implementing and studying other CO₂ sequestration and utilization systems to obtain time-resolved global warming metrics⁴⁹. In this process, it is important to underline that the lack of harmonization in assessment methods by LCA practitioners (dynamic vs. static approach, cradle-to-gate vs. cradle-to-grave, attributional vs. consequential, functional unit determination, low transparency in inventory modeling, standard variability) makes it difficult to benchmark across different methods for C sequestration. Consequently, evidencing their climate benefits, the trade-offs to each approach and any potentially occurring co-benefits quickly becomes cumbersome^{45,50}. Accordingly, international consortiums may be considered between LCA accounting methodologies that could interact with each other to ensure synergistic improvement

in accounting of sequestration potential of engineered and biosequestration CCS strategies alike.

Bio-sequestration strategies for CCS

Biological carbon sequestration (BCS) emerges from the complex interplay of biocatalytic processes conducted by microbes, plants and animals, including both direct (e.g., carbon fixation)⁵¹ and indirect (e.g., biomineralization)⁵² mechanisms. Consequently, in contrast with engineered CCS systems, the impact of BCS is less clearly delineated. Managed interventions on higher plants and microorganisms with the intent of removing CO₂ from the atmosphere is currently the focus. BCS has shown promise for its cost-effectiveness, ease of implementation and potential for long-term continuous sequestration by integrating seamlessly into local ecosystems as well as food production streams. Carbon can then take various forms, such as a richer microbiome, more plant biomass, a richer root network and soil organic matter (SOM), or more recalcitrant carbon forms, inclusive of oceanic sediments and peat, to name a few. Improving stored organic carbon (OC) should prioritize strategies according to effectiveness, permanence, associated ecosystem impact, and feasibility.

Ocean and wetland-based strategies

Many CDR approaches used in coastal, oceanic and wetland settings are based on augmenting the naturally occurring carbon sink processes within these ecosystems (**Figure 3**). Inorganic and organic marine carbon is found either as dissolved (DIC and DOC) or particulate (PIC and POC) forms. Of note, oceanic DIC levels are directly related to the metabolic activities of marine organisms (notably CO₂ uptake by photosynthetic processes of phytoplankton and bioproduction of carbonates by molluscs and coccolithophores). Although the vast majority of the oceanic pool is stored as DIC, marine sediments are still thought to contain approximately 2,322 Gt C in the top 1 m, and these stocks are also effective carbon sinks⁵³. Current estimates place ~85% of marine sediment stocks in the deep ocean⁵³, where they can be expected to remain for at least several thousand years⁵⁴, with similar estimates existing for DICs⁵⁵. The oceanic pool is thought to be undergoing an annual increase of 2.3-3.2 Gt C yr⁻¹, and is expected to further increase in the coming years^{56,57}. Anoxic marine sediments are also steadily accumulating⁵⁸. The high rate of sediment deposition, low oxygen levels, low hydraulic conductivity of sediment, and slower microbial decomposition rates all contribute to the burial and long-term storage of carbon in these sediments^{59,60}. The marine biological carbon pump (BCP), which drives the biological component of oceanic-atmospheric gas exchange and transportation of carbon to deep waters, has received much attention for its significant role in oceanic carbon cycling (**Figure 3a**). The BCP comprises 20 Gt C as particulate organic matter and is thought to account for sequestering 0.23-0.3 Gt C yr⁻¹⁶¹. The current accounting of carbon in oceanic stocks remains, however, fragmentary with no global totals reported on, e.g., suspended inorganic carbon, suspended organic carbon, or biomass carbon and relative contributions of biotic and abiotic processes driving oceanic carbon fluxes.

Of the emerging man-made strategies for marine BCS, oceanic fertilisation, alkalization and macroalgae cultivation have received the most attention. Ocean

fertilisation exploits the fact that approximately 25% of oceanic surface waters are thought to be high-nutrient, low-chlorophyll waters, where the growth of resident phytoplankton is hampered principally by micronutrient deficiencies due to precipitation of metal ions in their oxide form⁶². Ocean fertilisation approaches seek to sequester CO₂ in the deep ocean by inducing a bloom of phytoplankton with chelated Fe and soluble phosphate supplementation that subsequently sinks to deep waters (**Figure 3b**)⁶³. Although potentially inexpensive to implement in near-coastal areas, iron fertilisation carries side-effect risks. These include, for instance, disruption of ecosystem function (contributing to ocean deoxygenation due to increases in bacterial respiration of elevated organic carbon produced by phytoplankton⁶⁴) and pollution from the logistical dependency on maritime vessels.

Oceanic alkalization via the deliberate deposition of minerals (e.g., limestones and olivine) aims to enhance the solubility of CO₂ through the resulting pH shift, thereby raising its uptake into marine waters. This approach benefits from the sheer capacity of the oceanic pool and is thought to be capable of housing upwards of 800 Gt additional C by the year 2100⁶⁵. However, the long-term fate of the additional DIC is not entirely clear, though it has been suggested that combining alkalisation with carbon mineralisation may offer a means to circumvent this uncertainty in long-term storage⁶⁶. Importantly, there are considerable energy requirements for alkalisation at scale, and the resulting alterations in marine chemistry are not trivial. Such perturbations can be expected to alter the turbidity and trace metal levels, running the risk of disrupting biogeochemical cycles, biodiversity, and ecosystem services.

Macroalgae may sequester up to 0.166 Gt C yr⁻¹ via oceanic burial in coastal and deep-sea sediments^{67,68}, conceivably surpassing other coastal blue carbon systems in storage potential through cultivation initiatives. However, the impact of large-scale macroalgae cultivation (**Figure 3b**) on BCP equilibrium (due to nutrient sequestration and hypoxia in surrounding waters), warming (from decreased albedo), and mass transport of decaying organic matter to the seafloor sediments poses significant risks that must be carefully evaluated before implementation. Microalgae also contribute significantly to blue carbon (**Discussion S4**) by fixing CO₂ at rates up to 10 times greater than that of higher plants^{69,70} while producing valuable biomass products⁷¹ and may therefore find use in biotechnological applications that are carbon negative, though large-scale cultivation remains cost-limited. Chemoautotrophic processes in oceanic sediments are thought to naturally account for and long-term sequestration of carbon in stable mineral form (as much as 0.77 Gt C yr⁻¹)⁷². Despite these enticing features, chemoautotrophs in seafloor sediments have been challenging to adapt for BCS, owing to the limited accessibility to their environment and technical challenges associated with their manipulation.

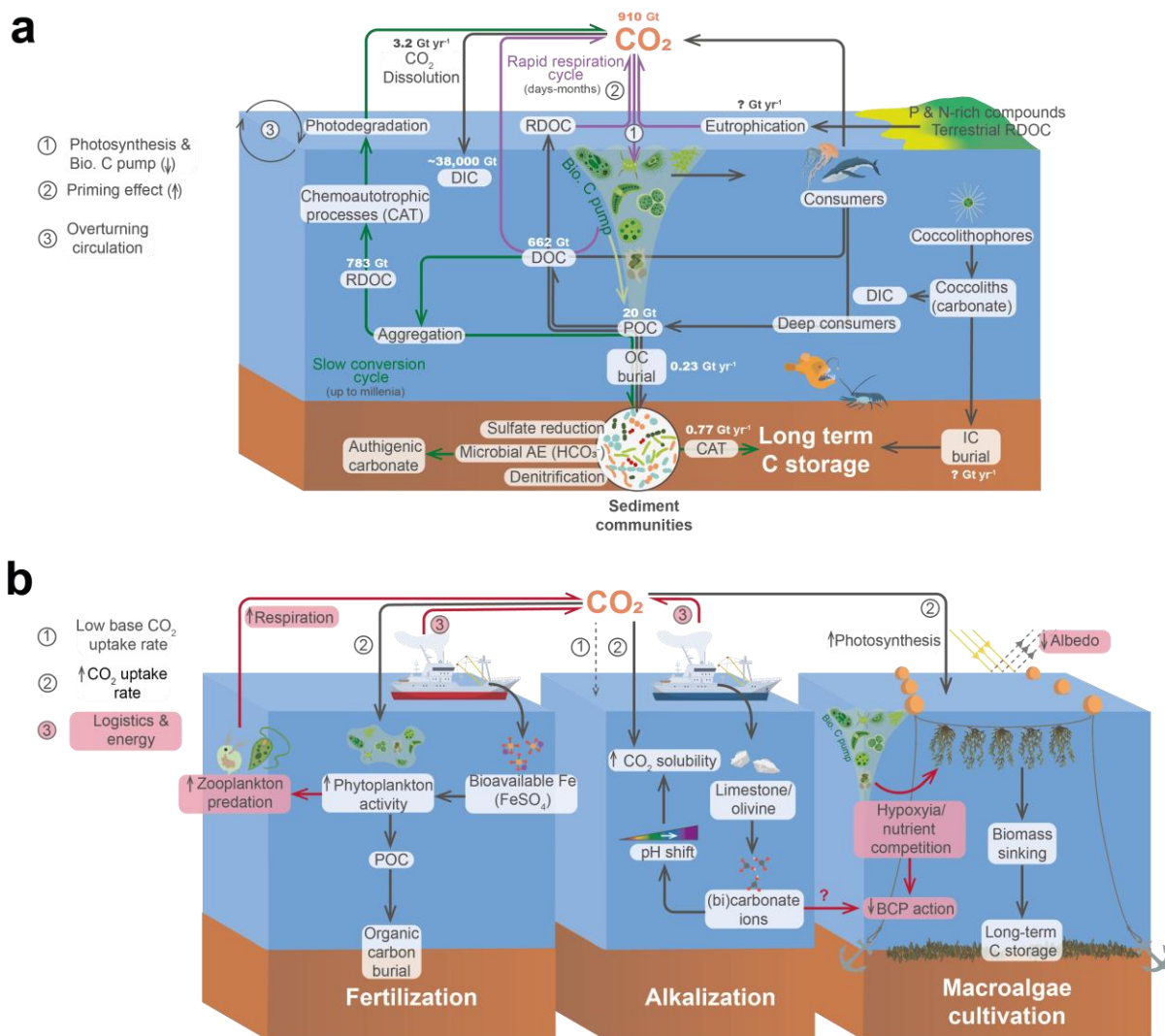


Figure 3 | Dynamics in oceanic carbon cycling and overview of oceanic CCS approaches. a - Ocean carbon cycling dynamics showing the principal mechanisms by which ocean carbon moves through the marine pools (DIC, DOC, RDOC, POC and biomass (comprising plankton and consumers)) and is either locked up for extended periods or released back into the atmospheric pool. Processes forming the rapid respiration cycle or slow conversion cycling are shown in purple and green, respectively. Pools with known magnitudes are presented (using mean values where ranges have been previously reported), and processes with known rates are also given; further information on data sources can be found in **Discussion S7**. Arrows within parenthesis indicate the direction of the process for clarity, where up indicates carbon release, and down indicates carbon uptake. **b** – Summaries of the major oceanic carbon capture and storage approaches (fertilisation, alkalization and macroalgae cultivation), dosing chemistry examples are presented for fertilisation (FeSO₄), and alkalization (limestone/olivine), though others also exist. Intended processes for each approach are shown in white boxes, and unintended side effects are indicated in red boxes. Where: P – phosphates, N – nitrogen, DOC – dissolved organic carbon, DIC – dissolved inorganic carbon, OC – organic carbon, IC – inorganic carbon, POC – particulate organic carbon, RDOC – recalcitrant DOC, AE – alkalinity enhancement, CAT – chemoautotrophic processes.

Blue carbon considerations also involve coastal systems such as halophilic flowering plants, namely tidal marshes, mangroves, seagrasses, coral reefs, and others^{73,74}. Among these, coastal flowering plants predominantly inhabit depositional shores as rhizophytes⁷⁵, where a portion of microbes can be retained as particulate organic

carbon in the root zone^{73,74}. Tidal salt marshes provide critical ecological services, including habitat for estuarine organisms, shoreline protection⁷⁶, and sequester an estimated 0.0048-0.0872 Gt C annually⁷³ attributed to their high sedimentation rates, elevated soil carbon content, and efficient burial of organic matter⁷⁷, while offering reduced methane emissions⁷⁸. Sediment brought in by tides, coupled with vegetation that acts as a baffle to enhance sediment deposition, supports the accretion process⁷⁹. Seagrass meadows have been highlighted to be able to store up to 19.9 Gt of organic carbon^{80,81}. Although seagrass meadows cover a mere 0.1% of the ocean floor, they contribute disproportionately to oceanic carbon burial, accounting for 10–18% of the total, with carbon accumulation rates ranging from 0.048 to 0.112 Gt of carbon annually^{58,81}. Coral reefs represent the most biologically diverse marine ecosystems, with an estimated annual global carbon sequestration of approximately 0.9 Gt C yr⁻¹. Their primary productivity spans from 0.3-5 Mg C ha⁻¹ yr⁻¹, indicating near-complete utilisation of CO₂ fixed during photosynthesis, primarily by mineralization⁸². It is, however, postulated that the precipitation of 1 mol of calcium carbonate (CaCO₃), even when buffered by seawater, results in the release of 0.6 mol of CO₂⁸³. CO₂ is utilised in the formation of CaCO₃ skeletons, suggesting that organic carbon metabolism can also serve as a net carbon sink whilst providing co-benefits of maintaining biodiversity and marine ecosystem services. Presently, coral restoration and creation efforts are not typically scoped for carbon sequestration. Integrating these initiatives within broader blue carbon strategies may realise as yet untapped synergies for climate mitigation and biodiversity conservation.

Beyond open marine waters, wetlands are effective for sequestering atmospheric CO₂ naturally^{84–86} as they possess the highest soil carbon density compared to forests and grasslands^{86,87}. Undisturbed wetlands can store carbon for hundreds to thousands of years; their protection, restoration, creation and the alteration of wetland components involved in organic matter decomposition are the main strategies for enhanced wetland C sequestration^{88–90}. For example, restoring drained European wetlands (~1.51 × 10⁸ hectares⁹¹) could sequester an average of 0.4 Mg C ha⁻¹ yr⁻¹⁹² (amounting to 0.06 Gt C yr⁻¹ across European wetlands). Biotechnological interventions for BCS in wetlands, such as manipulating soil microbes and vegetation, have been proposed to enhance carbon storage⁹³. Amendments such as biochar and humic acid application to wetland peat soils are also considered a viable strategy^{94,95}. Mangrove forests have emerged as notable wetland niches for their high-density C accumulation (owing to their capacity to drive C storage in both biomass and surrounding sediments). Indeed, at 738.9 Mg C ha⁻¹ on average, mangrove forests store more blue carbon per unit area than any other vegetative ecosystem, accounting for 10-15% of coastal sediment storage despite representing 0.5% of the global coastal area⁹⁶. Implementation of mangroves for BCS will, however, need to manage risks of methane emissions from anaerobic sediments in this niche (which can exceed several thousand μmol m⁻² h⁻¹)⁹⁷. Nevertheless, given their broad ecosystem services in providing habitats for a variety of marine wildlife, supporting nutrient cycling, and preventing coastal erosion, mangroves are an attractive niche for BCS applications. There, mangroves may find additional utility in hybrid nature-based solutions, for instance via restoration or cultivation for coastal and fishery catchment buffering or by integration with aquaculture for sustainable economic benefit provision^{98,99}.

Terrestrial strategies

Naturally occurring soil carbon uptake contributes between 0.4 - 1.2 Gt C yr⁻¹, or 5-15% of worldwide fossil-fuel emissions¹⁰⁰. The terrestrial sink, including the biotic pool, has been reported to absorb a net $\sim 1.4 \pm 0.7$ Gt C yr⁻¹, and models suggest this capacity may rise to approximately 5 Gt C per year by 2050¹⁰¹. The biotic pool provides the organic material that feeds the largest active source of atmospheric CO₂ from the pedologic pool, i.e., heterotrophic respiration (Rh), which drives the terrestrial pump (**Figure 4a** and **Discussion S5**). Rh releases an estimated net 35.4 ± 11.3 Gt C yr⁻¹^{102–104}, between 232% and 519% of the emissions released from anthropogenic CO₂ sources. The microbial activities underpinning Rh are governed by a variety of factors (such as solid or gaseous nutrient availability, temperature, and moisture) and remain an active area of research.

Terrestrial strategies for CCS centre on leveraging the biotic and pedologic carbon pools (**Discussion S6**). There, forests account for 40% of below-ground carbon storage¹⁰⁵. Methods to leverage forests for BCS included afforestation and reforestation. These carry some sequestration potential (up to 0.2 Gt C yr⁻¹ on repurposed grazing lands)¹⁰⁶, but they depend on site-specific factors and require substantial land and upfront investment^{107,108} which limits their scalability. Other land management practices (e.g., peatland restoration, paludiculture, and conservation agriculture) aim to harness the natural capacity of biological systems to capture and retain atmospheric CO₂ through photosynthetic processes and plant biomass accumulation, thereby stabilising soil carbon. Whereas soil amendments (such as biochar) augment existing stocks of, primarily, SOC with new carbon deposits in stable form. When considering land management practices, SOC plays a key role in the soil and ecosystem health, owing to the role of SOC in making soils more auspicious for biomass growth¹⁰⁹ (**Figure 4a**). Conservation Agriculture (CA) enhances SOC by minimal soil disturbance (no-tillage or minimum tillage), improving crop diversity (to enhance the variety of exudates and rooting depths) and permanent soil cover for increasing deep-seated plants¹¹⁰. However, CA is often conflated with reduced-tillage or no-tillage farming, leading to data variability and scepticism regarding its impact on SOC¹¹¹ (see also MRTs in **Figure 2**). Additionally, SOC outcomes can vary significantly depending on site-specific factors and management practices¹¹². The variability in SOC outcomes from CA highlights the need for further research to better quantify SOC gains and the associated increases in aboveground biomass resulting from improved land management.

Soil amendments (**Figure 4a**) such as compost, livestock manure, green mulches, and particularly biochar can enhance both microbial and plant biomass, nutrient and water retention while introducing degradation-resistant carbon that limits Rh-based turnover¹¹³. Among these, biochar (which retains >50% of the original biomass carbon¹¹⁴), offers one of the most promising mid to long-term CCS strategies, though its effects on SOC vary depending on application rate and soil context¹¹⁵.

Microbes are the fundamental regulators of SOM dynamics and nutrient availability (**Figure 4a,b**)^{116,117}. The management and manipulation of soil microbiomes is

therefore crucial. Biotechnologies centred on soil bioaugmentation, the introduction of pre-cultivated microbial cultures into the soil, have shown promise. Bioaugmentation enhances soil properties, increases plant growth yield, accelerates bioremediation and stabilises SOC (promoting aggregate and humified carbon formation)^{118,119}. However, consistency of outcome remains an ongoing research challenge owing to the complexity of soil dynamics and regional soil idiosyncrasies. The potential of specific microbes specifically designed to capture and sequester carbon has also been explored¹²⁰. This includes enhanced CO₂ sequestration via the production of urease- and carbonic anhydrase-producing microbes¹²¹, contribution to soil biomass through fungi's filamentous structures^{122,123}, and introduction of species capable of converting CO₂ to acetate via the Wood–Ljungdahl pathway¹²⁴. The production of carbonates from oxalate by oxalotrophic bacteria has presented another means for the accumulation of stable mineralised carbon forms like CaCO₃, which may be especially beneficial for facilitating CCS in soils cultivating oxalic acid secreting crops (such as Amaranth, Taro, and Spinach)^{125,126}. Indeed, we are at the point where the engineering and 3D printing of living materials capable of CO₂ mineralisation-coupled photosynthesis is now possible¹²⁷. Developments of this sort demonstrate that the continued efforts at the nexus of disciplines (e.g. ecology, synthetic biology, materials science) can be expected to yield novel biotechnologies that may contribute meaningfully at the local level as modular tools within a holistic and integrated framework for sustainable habitat integration.

The balance between carbon forms and microbial metabolism (where aerobic pathways favour respiration and anaerobic conditions promote methanogenesis) significantly influences the re-emission of organic carbon. Ideally, captured atmospheric carbon could be stabilised for long-term storage via the intervention of manufactured systems (e.g., pyrolysis), natural carbon pumps or by biotransformation. However, for most carbon-saturated soils, for example, in areas with rich soils such as forests and wetlands (Figure 4b), further introduction of organic carbon typically leads to re-emission, highlighting the need to consider the localised starting C content of soils or ecosystems where BCS strategies are implemented.

Leveraging agro-climatic zones for favourable biological carbon soil equilibria

For carbon-saturated soils (**Figure 4b**), the equilibrium between storage and respiration of carbons is usually stable and further carbon storage involve transforming OC into more stable carbon forms. In contrast, soils that initially have a low carbon content may favour carbon storage from OC given that few organisms are present for re-emission (**Figure 4a, right**). Among the C poor regions, the MENA region, characterised by arid and semi-arid climates, faces distinct challenges for BCS due to water scarcity and land degradation. However, innovative land management practices could offer outstanding BCS benefits. Agroforestry, integrating drought-resistant tree species with crops, might enhance carbon sequestration and provide additional benefits such as improved microclimates and soil stabilisation, particularly in areas with supportive socio-economic frameworks for sustainable agriculture. Coastal areas in the MENA region, particularly along the Red Sea and the Mediterranean, hold potential for blue carbon sequestration through the restoration of mangroves and salt marshes.

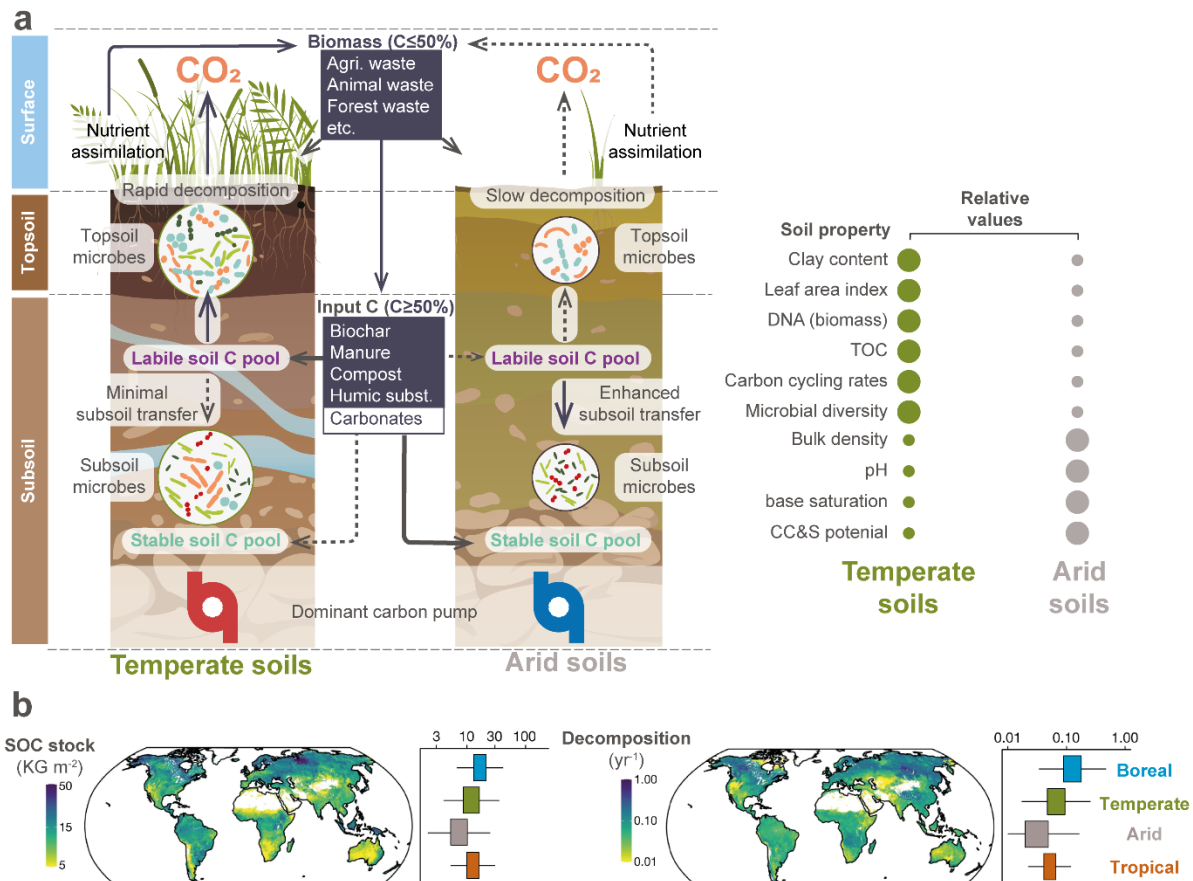


Figure 4 | Carbon saturated (at maximum carbon content equilibrium) and carbon-poor soils (far from maximum equilibrium) and their global predominance. a - responses of temperate and arid soils to input carbon are distinct owing to their respective carbon cycling dynamics which are emergent properties of the local soil-associated biological systems and physicochemical soil properties. Percentages shown refer to the typical carbon content in the biomass and processed input carbon. The dominant carbon pump in temperate soils (red) is microbial (characterized by high carbon turnover, heterotrophic respiration and organics and a correspondingly low capacity for additional carbon input). The dominant carbon pump in arid soil (blue) is the mineral carbon pump (characterized by low carbon turnover mineral adsorption, mineral occlusion and aggregation and a correspondingly higher capacity for additional carbon input). Further details on the mechanisms underpinning these systems are provided in **Figure S1**. Where: Agri. Waste - agricultural waste biomass, Humic subst. - processed humic substances, TOC - total organic carbon and CC&S - carbon capture & storage. **b** - Global heatmaps showing soil organic carbon (SOC) stocks and relative extents of soil decomposition. Adjacent box plots show the breakdown of these values for the four principal soil types (boreal, temperate, tropical, and arid), demonstrating the favourable SOC stocks and decomposition rates in arid soils. Plots adapted from Tao et al 2023¹²⁸.

SOC stability is a critical component that is governed through the complex interplay of biological, chemical, and physical factors¹²⁹. The importance of targeting carbon-poor (e.g. arid) areas for a successful BCS implementation is becoming increasingly recognized. For example, studies on the priming effect (PE) describing the rate of organic matter mineralisation, and biochar's influence on this process^{130,131}, have shown predominantly negative PE values (a reduction in SOC decomposition), except in sandy soils (20.8% positive PE) and other poorly fertile soils⁴¹. Furthermore, afforestation has greater potential as a carbon sink on barren lands with low initial SOC stocks compared to croplands or pasturelands. In the latter, fresh SOC inputs

(e.g. from tree litter) can recharge the existing SOC pool and prompt the PE, accelerating SOC loss^{132,133}. Whereas, in carbon-poor soils, microbial activity is limited due to the lack of bioavailable carbon. There, any new addition tends to exceed the decomposition rate and causes a commensurate increase in SOC stock. Indeed, microbial activities (such as Rh) in arid soils are further restricted by the lack of available water, promoting the retention time for newly added organic matter. However, such regions are also prone to accelerated erosion from wind exposure, which can strip topsoil and undermine SOC accumulation in the long term. Water scarcity must also be carefully considered as it constrains biomass production and microbial establishment, limiting the natural capacity for sustained SOC buildup and thereby hindering the establishment of ecosystem services. Notably, an estimated 36% of tropical land available for afforestation receives only 40% of the necessary water by rainfall, exacerbating local water scarcity¹³⁴. The synergy between BCS and water supply will therefore be critical in realizing gains for the most out-of-equilibrium areas where carbon-poor soils are prevalent.

Shaping of atmospheric composition by biological systems

While equilibrium processes govern carbon storage at local scales, it is the cumulative activities of biological systems across environments that have ultimately shaped the Earth's atmospheric composition. Over geological timescales, biological activity has profoundly altered the atmosphere, driving major shifts in global carbon cycling (**Figure 5**)^{135–137}. During the Archean eon, N₂, CO₂ and CH₄, were two to four orders of magnitude higher than today¹³⁵. This composition lasted until the introduction of oxygen, likely through microbial photosynthesis¹³⁸. Indeed, the latest model of CO₂ over the past 485 Ma is consistent with the marked decline in atmospheric CO₂ following the emergence of land plant life¹³⁹. While atmospheric CO₂ has a moderately strong correlation with global mean surface temperature (GMST) across these timespans ($r \approx 0.57$), other climatic factors are clearly also at work in shaping the full climatic signal (**Figure 5a**). Records for other major GHGs (i.e., CH₄ and NO₂) are less complete, owing to their respective instability and lack of geological proxies, though ice-core data are available for the past 800 Ka^{140,141}. Across this timespan, a coordinated rise and fall in these GHGs can be seen with a cyclic periodicity of ~100 Ka (**Figure 5b**). This periodicity corresponds to orbital eccentricity, a key aspect of Milankovich cycles, that perturbs climate via insolation-driven orbital forcing¹⁴². Orbital forcing works globally on GMST, as an external climate driver, influencing climate feedback systems including ice cover, ocean circulation and biogeochemical cycles. This interplay of orbital dynamics and GHGs highlights the importance of developing our understanding of how Earth's systems, both biological and otherwise, modulate responses to astronomic forcing.

Microbiological metabolic networks are central to natural GHG cycling. Key microbial players include: **1)** Syntrophic fermenters, which generate CO₂ and provide the H₂ and acetate for methanogens, which produce CH₄, interfacing with the nitrogen cycle. **2)** Ammonia producers, which inhibit aerobic methanotrophs that would otherwise oxidise CH₄ into CO₂. **3)** Denitrifying methane oxidisers, which consume CH₄ to produce N₂ and CO₂. These microbial systems are complex and deeply intertwined, and yet, are often studied in isolation¹⁴³, which limits our capacity to model their role

in atmospheric dynamics or manage their cumulative impacts. Despite their established contemporary impacts and suspected historical roles in shaping atmospheric composition, biological systems remain underutilised in modern CCS strategies, where the focus remains on engineered solutions^{144–146}. This is partly due to the relative complexity of biological systems and persistent knowledge gaps surrounding their workings. Notably, the nature of below-ground microbiome responses to dynamics in atmospheric GHGs is poorly understood and challenging to predict¹⁴⁷. This is further complicated by uncertainties in the coupling of CO₂, CH₄ and NO₂ cycles (**Figure 5**), where, for example, methane emissions offset CO₂ storage in coastal ecosystems due to methanogenic activities in sediment layers¹⁴⁵. These microbial complexities highlight the need for a cautious and systems-based approach to BCS¹⁴³. Nevertheless, the engineering of biomass and biocatalysis may represent one of the most sustainable avenues for global engineering of atmospheric, soil, and water body compositions while strategically providing for society's needs.

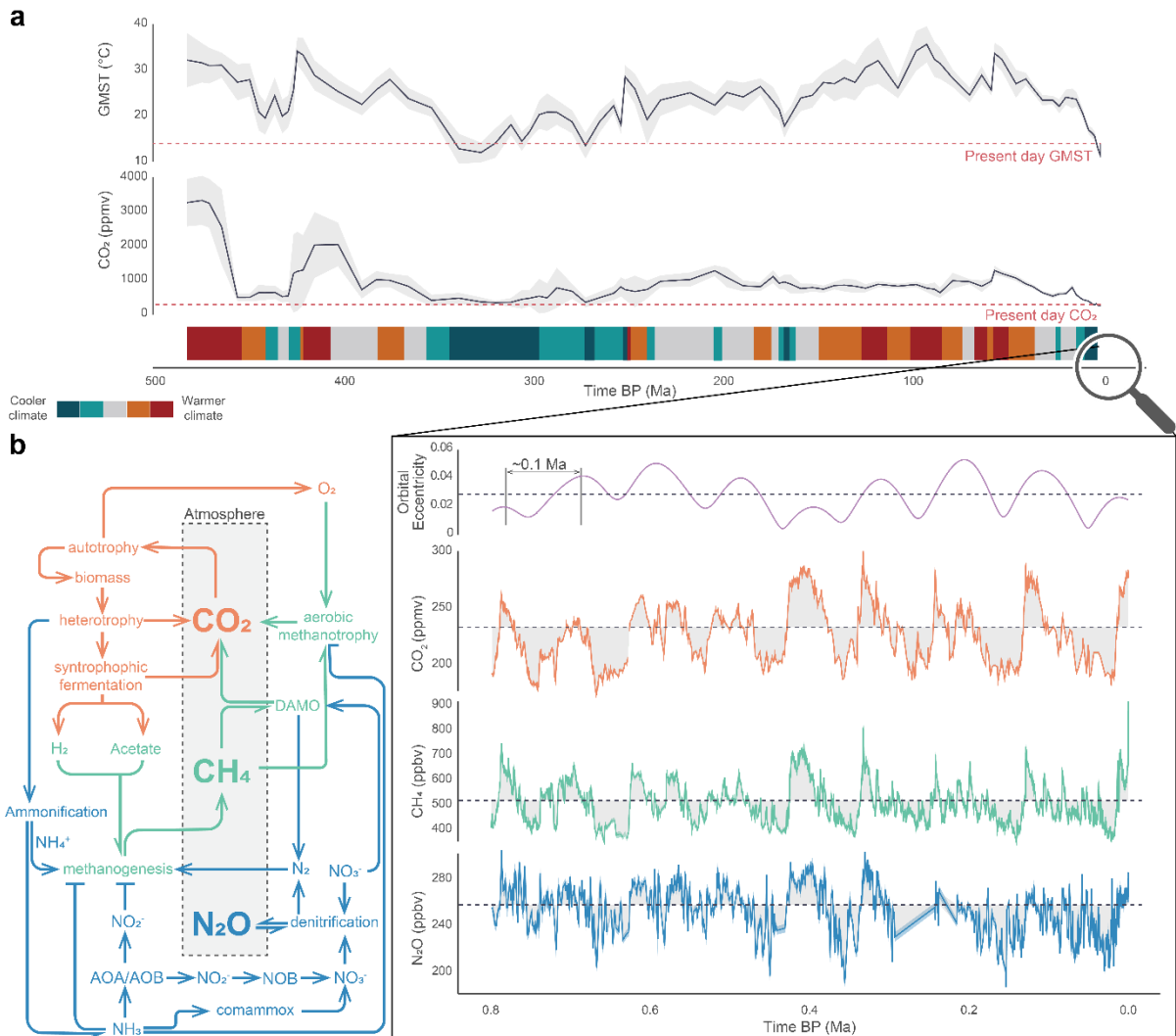


Figure 5 | Highlight of geological impact of biological systems on compositional dynamics of the atmosphere via the balanced interplay of interlinked biogeochemical cycles. a - Global mean surface temperature (GMST) together with corresponding CO₂ levels over the past 485 Ma. The coloured bar highlights qualitative climate conditions (from cooler to warmer). **b -** the profiles of Earth's elliptical orbit eccentricity profile alongside major biogenically-produced GHGs (CO₂, CH₄, N₂O) over

the 800 Ka history available from the Vostok and EPICA Dome C ice core records. Cycles in these gases (with a periodicity of ~100 Ka) are highlighted by the cartoon above, the medians are presented as dashed lines to convey a sense of the theoretical equilibria for these GHGs, the fluctuations about which are highlighted by the grey-shaded areas between the data and the median lines. The transparent shading around the data (coloured by gas), depicts the uncertainty range ($\pm 1 \sigma$). **c** - are the biogeochemical cycles of these GHGs, emphasising the connections between them (coloured by gas, as in the chart). Where: DAMO - denitrifying anaerobic methane oxidation; AOA/AOB - ammonia-oxidising archaea/bacteria; NOB - nitrite-oxidising bacteria. Details on data sources can be found in **Discussion S7**.

Conclusions

The impact of biological systems on carbon cycling far exceeds that of any current or emerging manufactured CCS system. For example, shifting the balance of carbon exchange associated with natural systems by 0.1% would surpass the total annual impact of current engineered CCS successes to date. Biological processes offer dynamic and often scalable means of capturing carbon across ecosystems. These global processes can be leveraged to inform BCS strategies that enhance carbon capture while delivering ecosystem services, such as food, materials, and other bioproducts. As such, BCS strategies are well-positioned to be embedded in future strategies for food and water security, as well as ecosystem management.

A deeper understanding of how biosystems interact with carbon pools is therefore crucial for assessing the spatial and temporal dynamics of carbon cycling and informing carbon management interventions that are both effective and ecologically sound. In this context, evaluating the interactions between, and the relative impacts of, emerging BCS will be increasingly critical.

The carbon retention timescales of BCS are broad, from decades (as in ocean fertilisation or soil amendments) to millennia (as in peatland restoration and biomineralisation), which offer more permanence. In contrast, abiotic geological storage solutions, such as carbon injection into depleted hydrocarbon reservoirs and saline aquifers, provide near-indefinite sequestration¹¹ with lower biological interference, but are often energy-intensive and lack ecosystem co-benefits. Nevertheless, both current and future approaches may find their place as part of an integrated, multi-scale, and systems approach to carbon management adopted across diverse settings. However, the contrast between biological and abiotic sequestration pathways highlights a key consideration for integrated carbon management wherein biological processes are leveraged where short- to medium-term sequestration is needed, while ensuring permanent storage options are in place for long-term stability.

A sustainable carbon management framework will require advances in our understanding of biological dynamics, interfaces between biological systems as well as in our ability to monitor carbon transformations and co-emitted gases (CH₄ and N₂O) that reflect microbial activities and co-influence climate. This includes further investment in robust carbon accounting dynamics, improved costing frameworks, and a reduced reliance on uncertain models will be critical. Tailoring strategies for regionally specific settings is equally important, as the variability in “carbon-farming”

and certain other biogenic approaches has shown. Biotechnologies and the creation of circular economies certainly offer promising potential for carbon-based climate management. Nevertheless, large-scale biogeochemical interventions carry unprecedented scope, socioeconomic and ecological implications. The necessary degree of predictive control required for this to be done safely, therefore, requires that restraint, adaptability, and grounding in ongoing observation must be the cornerstones of any responsible carbon strategy. In this way, we can ensure that we are building toward a resilient, regenerative, and sustainable carbon management framework that supports both climate stability and continued ecosystem integrity into the future. Looking ahead, our ability to manage biosphere-atmosphere interactions may prove critical for long-term resilience, not only for climate mitigation, but also in buffering future climate perturbations.

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Declaration of interests

We, the authors and our immediate family members, have no financial interests to declare.

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Supplementary Information

Biological Dynamics for Sustainable Carbon Sequestration

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Climate change mitigation; Soil carbon storage; Blue carbon ecosystems; Carbon storage lifetimes; Ecosystem services; Biological carbon pump; Nature-based climate solutions

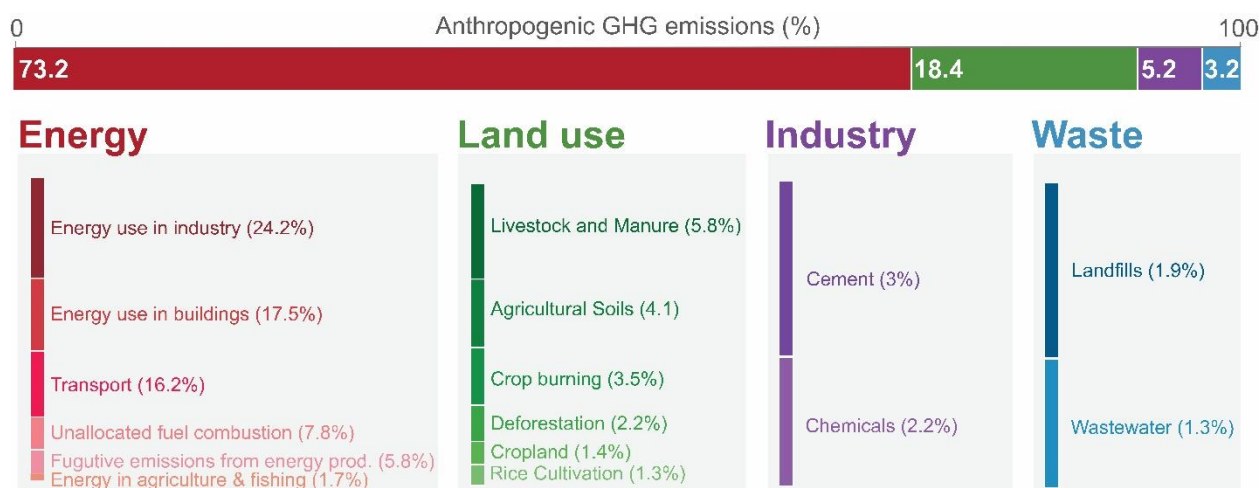


Figure S1 | Major landmarks of the global carbon (C) cycle and opportunities for intervention Summary overview of anthropogenic emissions by sector (as of 2021) as a percentage of the total anthropogenic release. The major subgroups within each sector are presented in the same colour series as normalised bars (with percentages reflecting the associated emissions of the total, such that the subgroup total equates to the value given for the sector). Sector GHG emissions data was taken from¹.

Discussion S1. Carbon costing

Carbon costing is necessary to ensure that selected management practices deliver genuine net benefits by balancing sequestration effectiveness against economic and ecological impacts, thereby guiding policy decisions and management practices toward long-term sustainability. The distribution of costs across time (i.e., upfront, monitoring and maintenance costs) needs to be factored in, particularly when evaluating engineered versus nature-based solutions. Additionally, the discount rates, commonly used for carbon costing, steer policymaking and investments in CCS (with lower rates encouraging stronger regulatory action and prioritising long-term benefits). These discount rates thereby hold strong influence over comparative analyses of engineered (where short-term capture gains may be high) and nature-based (where payoffs may be delayed due to biological dependencies such as biomass growth) solutions. Importantly, carbon costing discounts are not trivial to calculate, are set somewhat arbitrarily and do not account for co-benefits (e.g., environmental) of the capture method. Indeed, small variations can significantly alter the results of cost-benefit analyses, and caution is therefore warranted in their use. Consequently, there is no clear consensus on where they should be set and the degree to which they should be implemented, though they are generally on an upward trend².

In any event, levelized costs of carbon capture are a critical determinant of scalability for CDR (S. Table 1). Levelized costs for oceanic sequestration approaches, for example, vary between 7-17K+ USD per tCO₂ mitigated (S. Table 1), which equates to 25.66B – 62.33T USD per Gt C captured (or 0.249T - 605T USD to mitigate the current annual production of 9.7 Gt in anthropogenic CO₂). As with most CDR approaches, current costs are prohibitively high for full scale-up, requiring reductions, in some cases by several orders of magnitude, to approach economic feasibility.

Standardized and effective guidelines for the measurement, monitoring, reporting and verification (MMRV) of CDR programs are still needed for accurate carbon accounting³. Universally accepted MMRV guidelines would enable a more holistic consideration of new CDR programs by accounting for unintentional ecological, socioeconomic, and ethical impacts. A key challenge in the construction of MMRV guidelines will be in balancing the costs of acquiring adequate field measurements (to avoid overdependency on predictive models, which may be based on an incomplete view of the CDR setting) whilst ensuring a favourable cost-benefit result.

Supplementary Table 1 | Carbon costing for a selection of biological and engineered CCS approaches

Technology	Costing (USD per tCO ₂)	Sources
Direct Air Capture	80-1,133	4
Bioenergy with CCS	32-321	5,6
Industrial CCS	40-110	7
Enhanced weathering	60-200	8
Ocean fertilization	7-4,691	9-11
Biomass slurry fracture injection	~95	12
Kelp aquaculture CCS	1,257-17K	13
CCS for power plants	22-56	14,15
Biochar	30-139	16
Afforestation	220-1,165	17,18
Ocean alkalization	130-295	19

Discussion S2. Biologically-derived materials for a circular carbon economy

Nearly 106.1 Gt of processed material enter the global economy annually. Carbon constitutes an essential component of many of these materials (e.g. plastics), and its demand is expected to increase at a compound annual growth rate of 2.5%, rising from 0.55 Gt in 2020 (comprising 65% polymers and 35% organic chemicals), to 1.150 Gt in 2050²⁰. As a result, modern materials are a relevant source of carbon emissions. However, these materials could become carbon sinks upon model transformation by providing long-term C storage and avoiding carbon-intensive production processes²¹. In this context, while full decarbonization is impractical due to the C-based nature of materials, defossilization strategies that aim to reduce atmospheric fossil-derived carbon (below-ground) levels by progressively phasing it out with alternative feedstocks are more viable. However, certain applications are expected to continue to rely on fossil-derived carbon until viable substitutes are available at scale. Prioritizing sectors where the substitution of fossil-derived carbon is technologically and economically feasible offers promising inroads to prevent the influx of fossil carbon into above-ground systems (atmosphere, technosphere, biosphere), thereby curtailing increases in atmospheric carbon²².

Unlike most materials currently in use, which originate from fossil-derived carbon sourced from below ground and follow a linear extraction-processing-utilization-waste model, biologically-derived materials contain renewable biogenic carbon that, when kept in use within the technosphere, avoids its release into the atmosphere. Fostering biological material utilization follows the IPCC recommendations and reduces the need for complex carbon capture and storage systems²³. Wood (and derivatives like cellulose and lignin), chitin and silk are the most representative example of materials containing biogenic carbon. For every kilogram of embedded carbon in these materials, 3.66 kg of CO₂ are prevented from being released into the atmosphere (considering CO₂ and C molecular weight as 44 and 12 g·mol⁻¹, respectively)²⁴. As biologically-sourced materials typically contain between 36-64 wt% C (depending on species and processing), 1 kg of these materials locks up between 1.30 and 2.33 kg

CO₂ equivalents, providing potential for carbon sequestration in long-term applications²⁵. Furthermore, impressive functionalities can be obtained from woody materials, enabling the replacement of existing non-renewable materials in high-tech applications as well²⁶.

As biologically-derived materials could enable mild processing routes for fabrication, reduced *cradle-to-grave* (considering the entire lifecycle) CO₂ footprint than fossil counterparts are generally realized (Zuiderveen et al. recently reported 45% lower emissions with a 95% confidence interval)²⁷. Phasing out fossil-derived materials in favor of renewable C-containing materials (e.g., secondary materials and biomass utilization) would also increase the circularity of the global economy, which has steadily declined from 9.1 % (first Circularity Gap Report in 2018) to 6.9 % (Circularity Gap Report 2025)²⁸. Recent estimates suggest that it is feasible to cover 20% of the overall carbon demand of the chemical and derived material sector with agricultural, woody, and biowaste biomass by the year 2050²⁹. Therefore, the defossilization of carbon embedded in the materials and chemical industries should be accompanied by the recycling of existing C-containing materials (to keep them within the technosphere), and the implementation of CCU strategies for the materials sector (i.e., copolymerization of CO₂ with epoxides to produce polycarbonates)³⁰, to establish a truly circular carbon economy³¹.

Discussion S3. Carbon retention lifetime distributions

Illustrative log-normal carbon retention lifetime distributions were generated from minimum, maximum, and mean retention times and their standard deviations from data reported previously (S. Table 2). For instances where generalised terms for timeframe were given (e.g., *months*, *decades*), these were handled in a standardised manner using values as laid out in S. Table 3. In the case of the afforestation/reforestation category, minimum retention values were taken as the reported time in years to net C storage, whereas mean and standard deviation in this subgroup were based on an assortment of reported lifetimes of deciduous and coniferous trees of the sort used in such initiatives. In the case of algae cultivation, fuel and feedstock applications were factored in assuming a *weeks* to *months* timeframe for production to utilisation, and this subcategory also includes open ocean macroalgae cultivation and sinking data. Paludiculture (biomass cultivation on wetted peatlands) is an emerging CCS strategy, one that is potentially promising as a nature-based CDR solution. However, there presently appears to be insufficient C retention data to place it alongside the other strategies. Studies measuring the reduction in peatland emissions and the accumulation of biomass are therefore desirable. Log normal distributions were calculated with the *dlnorm* function using the retention time data (extended to 1/10th the minimum retention time and 10× the max retention time, for smooth tails), $\mu \log$ (the adjusted mean of the log transformed retention times, adjusted to account for skews in the distributions by subtracting half of the variance in $\sigma \log$), and $\sigma \log$ (the log-space standard deviations) in RStudio³². The script to reproduce the plots is available in **Supplementary Information 2**.

Supplementary Table 2 | Carbon Retention Time Data for CCS

Group	Subgroup	Min RT ¹	Max RT ¹	Mean RT ¹	St. Dev. RT ¹	Sources
Waste Materials	Landfill sequestration ²	0.082	100	34.216	41.974	33–36
	Waste biomass utilization (biochar) ²	10	999	303	466.325	37–39
	Waste biomass utilization (durable goods)	0.082	99	12.91	23.966	40–47
	Afforestation/reforestation ³	13	1,000	532.25	291.091	48–51
Plant Biomass	Agroforestry	1	50	28	21.608	52,53
	Algae cultivation ²⁴	0.019	9,999	2,230	4,361.672	54
	Blue carbon ecosystems	21	3,350	1,293	1,797.979	55
Soils	Conservation Agriculture	15	117	47.375	22.295	56,57
	Biochar Application	0.296	1,554	857	600.248	58,59
	Peatland Restoration	5,000	11,700	9,575	3,084	60–63
	Paludiculture	<i>Data not found</i>				
Man-made Products	Wood construction materials	25	100	48.75	34.731	64–66
	Bio-based materials ²⁵	0.005	10	0.5	2.499	67
	Bioconcrete	20	280	133.333	98.995	68
	Biomining ²	100	1,000,000	500,050	707,036	69
Oceanic Storage	Packaging materials ²	0.019	3	0.8664	1.254	70,71
	CCU fuel products ²⁶	0.019	1.89	0.743	0.884	-
	Ocean fertilization	0.986	100	37.8	50.046	72
	Alkalinity enhancement ²	50	1,000	441	399	55,73
	Carbonate mineral formation ²	100,000	999,999	550,000	636,395	73
Geological Storage	Depleted HC reservoirs ²	20	99,999	18,222	33,226	74–77
	Unmineable Coal Seams ²	550	10,000	6,512	4,505	73,78
	Saline Aquifers ²	550,000	9,999,999	3,025,000	3,941,034	73,79

¹ Retention times given in years.

² Literature includes generalizations, in these cases data is supported using figures from S. Table 3 below.

³ Data based on lifetimes of trees housing stored carbon (as described in Section 2 above)

⁴ Algae cultivation for fuel and feedstocks factored in assuming a weeks to months timeframe for production to utilization and includes open ocean macroalgae cultivation and sinking data.

⁵ Short-lived biomaterials such as bioplastics, natural fibres, adhesives, and fungal byproducts.

⁶ Assuming fuel storage time of days to months and immediate release upon use.

Supplementary Table 3 | Figures Used for RT Distributions when Handling Generalisations.

Handling generalisations	Min (years)	Max (years)	Mean (years)	SD (years)
<i>Days</i>	0.003	0.036	0.019	0.023
<i>Weeks</i>	0.019	0.134	0.077	0.081
<i>Months</i>	0.082	1.890	0.986	1.279
<i>Years</i>	1	19	10	12.72792206
<i>Decades</i>	10	99	55	63
<i>Centuries</i>	100	999	550	636
<i>Thousands of years/Millenia</i>	1000	9,999	5,500	6,363
<i>Tens of thousands of years</i>	10,000	99,999	55,000	63,639
<i>Hundreds of thousands of years</i>	100,000	999,999	550,000	636,395
<i>Millions of years</i>	1,000,000	9,999,999	5,500,000	6,363,960

Taking waste materials as an example, in landfills, microbes play a crucial role in carbon cycling, converting biodegradable waste into methane (CH₄), carbon dioxide (CO₂), and microbial biomass via both aerobic and anaerobic pathways. In these environments, Rh, methanogenesis and denitrification, are key pathways for carbon transformation⁸⁰. Saprotrophic activities of fungi are also important determinants of carbon cycling and can accelerate the degradation of recalcitrant substrates such as textiles and plastics^{81,82}. These processes vary with landfill depth and age, affecting organic matter degradation and carbon sequestration lifetimes⁸³. The retention times is affected by complex drivers, e.g. both the composition and nature of the sequestered carbon (e.g. lignin content, bioavailability, chemical treatments etc) and broader conditions of storage. Generally speaking, rapidly degrading products, such as food items, typically release their stored carbon within a year, whereas newspapers and wood items can last for decades or more. Carbon stored as a utilisation of waste biomass (engineered wood, biochar, textiles packing materials and bioplastics), is typically retained for longer (months to decades). However, retention times can be stunted when the biomass is not produced with sequestration applications in mind, as has been seen for waste biochars created from suboptimal pyrolysis and feedstocks, wherein mean retention times (MRTs) can be reduced down to decades versus 10²-10³ years of biochar created for CCS applications⁸⁴.

Discussion S4. Algae and blue carbon storage

Microalgae and macroalgae are also important drivers of blue carbon (**Figure 3**). Microalgae are estimated to produce approximately half of the atmospheric oxygen while simultaneously utilising CO₂ for their growth⁸⁵. With an airflow containing 4% CO₂, microalgae can fix carbon at a rate of 0.0146 Mg C ha⁻¹ /day while growing at 0.0302 Mg⁸⁶, up to 10x that of higher plants⁸⁷. CO₂ bio-fixation by algae also yields valuable biomass products. These include proteins, fatty acids, vitamins, minerals,

pigments, dietary supplements for humans and animals, and other bioactive compounds⁸⁸. However, large-scale cultivation projects of microalgae still remain costly, primarily stemming from their comparatively low productivity on a large scale. Initial calculations indicate that macroalgae globally facilitate an export of 0.679 Gt C yr⁻¹^{89,90}. Specifically, an estimated 0.014 Gt C yr⁻¹ is sequestered in coastal sediments, and 0.152 Gt C yr⁻¹ that could potentially be sequestered in the deep sea⁸⁹. It is reported that a part of macroalgae-derived POC is buried in shelf sediments, and a considerably higher amount is sequestered in the deep ocean⁹¹. These calculations imply that macroalgae might function as a larger sink for carbon than seagrass, tidal marshes, and mangroves combined. Nevertheless, macroalgae predominantly inhabit eroding shores, limiting the storage of particulate organic carbon near their production sites.

Discussion S5. Soil dynamics and the microbial carbon pump

Microbes have long been recognised for their vital role in soil carbon cycling. Recent work has led to the conceptualisation of the soil microbial carbon pump (MCP) and carbon use efficiency frameworks^{92,93} (**Figure S2**), which essentially serves as a terrestrial counterpart to the more established marine biological carbon pump. The soil MCP describes the microbial processes that collectively convert input soil organics into persistent SOC, thereby contributing to long-term carbon sequestration. This function of the MCP is achieved through the combined effects of anabolic assimilation processes (wherein organics are taken up and converted into microbial biomass)⁹², enzyme secretion and exogenous enzymatic carbon transformation (whereby organic matter is converted into metabolic byproducts and biopolymers, as in exudates and biofilms)^{94,95}, the accumulation of necromass (the cellular debris of dead cells)⁹⁶ and the production of recalcitrant carbon (whereby microbially-derived products stabilize carbon via their carbon interactions, often through aggregate formation)⁹⁷. Given the multitude of interfacial dimensions these processes make with other soil components (root systems, soil fauna and the soil matrix), they are complexly interwoven with soil settings. The MCP is therefore influenced by soil chemistry, soil physical conditions, and events above ground (especially those affecting topside flora⁹⁸), and yet highly adaptive to soil perturbations^{93,99–101}. Indeed, it is becoming increasingly evident that multimodal analyses (i.e., supporting physicochemical measures with integrative omics microbial measures taken across systems levels, e.g. metabolomics, metaproteomics, metatranscriptomics and metagenomics) are needed to properly capture microbial dynamics in soil systems^{102,103}.

Microbial communities are distinct across soil compartments and so can vary substantially along the depth¹⁰⁴ and oxygen¹⁰⁵ gradients. This can have pronounced implications for carbon cycling, where topsoils are typically associated with high turnover carbon release from the labile carbon pool¹⁰⁶, most notably via Rh and, in anaerobic settings, methanogenesis, processes which are interlinked with nitrogen cycling (thereby influencing N₂O emission)¹⁰⁷. Conversely, the subsoil is associated with low carbon turnover and long-term carbon storage through so-called *dark fixation* (principally driven via the chemoautotrophic actions of nitrifying bacteria, and sulphur or iron-oxidizing prokaryotes)^{108–111}. Consequently, the soil MCP has a direct bearing on soil carbon fates and soil microbes should therefore be a key consideration in site

selection for soil carbon sequestration initiatives (particularly those based on soil amendments).

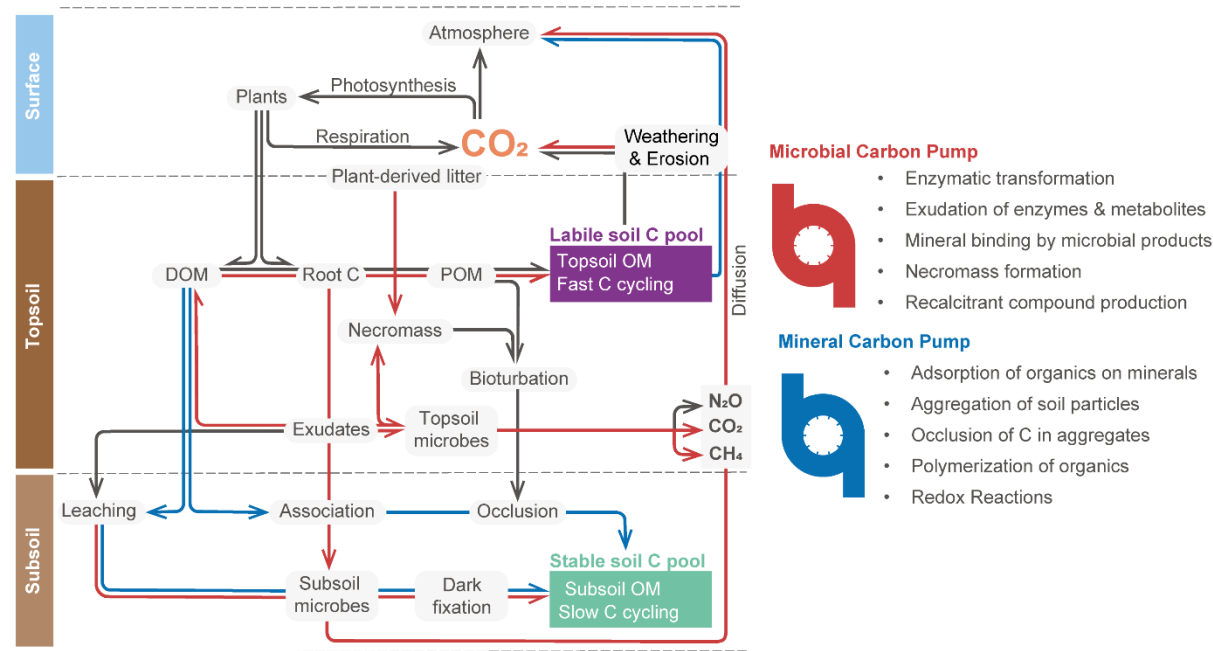


Figure S2 | Soil carbon dynamics are driven by the microbial or mineral carbon pumps. Soil processes are delineated (dashed lines) across soil compartments. Process lines are coloured by carbon pump type (where those in red are driven by the microbial carbon pump and those in blue by the mineral counterpart) remaining processes are in black. On the right, the key features for each pump are summarized in bullet form.

Discussion S6. Forestation, land use and soil capacity

Considering the fact that forests store 40% of all below-ground carbon¹¹², afforestation, the expansion of forested areas, has been put forward as a CCS approach. Since the late 20th century, global afforestation efforts have increased planted forest areas by approximately 105 million hectares¹¹³. Afforestation and reforestation, while highly effective at sequestering carbon, require larger land areas and significant upfront investment. CA, afforestation and reforestation are influenced by factors such as tree species, land use history, stand age, climate, and soil type^{114–117}. Studies indicate that during the initial 20 years of reforestation, aboveground biomass increases at a rate of 6.2 Mg ha⁻¹ yr⁻¹, which slows to 2.9 Mg ha⁻¹ yr⁻¹ over the next 80 years of regrowth. Globally, there are an estimated 3.5 billion to 3.7 billion hectares of grazing land available^{113,118}, potentially enabling sequestration of up to 0.2 Gt C yr⁻¹¹¹⁹, to the extent that ethical and socioeconomic implications are reasonable. Although the optimal approach will vary based on regional factors and the specific dynamics of each ecosystem, effective forest management strategies and adaptive grazing practices (such as high density-short duration grazing)¹²⁰, establishing grazing exclusion zones¹²¹, proactive soil intervention¹²² and coupling strategic grazing plans with forage species diversification efforts¹²³ can significantly support both carbon sequestration and ecosystem services.

SOC amendment is an alternative strategy that has the potential to increase soil microbial biomass^{92,124}, water retention¹²⁵, and introduce recalcitrant C. Livestock

manure, compost, biochar, and green mulches are some of the common soil amendments (**Figure 4a**). Additionally, ongoing research on other possible innovative amendments include treated municipal sewage sludge and aluminium or steel slags^{126–129}. Biochar amendments also provide stable carbon forms that are resistant to microbial decomposition, limiting Rh-based turnover and standing out as one of the most promising soil amendment approaches for CCS. Biochar is produced through the pyrolysis of organic materials, transforming OC into stable refractory products with a half-life ranging from millennial to centennial timescales. Biochar typically retains 50% C from its original biomass, and up to 100% in idealized conditions, compared to combustion and composting, which retain only around 3% and 10–20%, respectively^{130,131}. Furthermore, optimal C sequestration occurs with H/C ratios between 0.38 and 0.44, which is achieved at the highest treatment temperatures (HTT) of 500–550°C across diverse feedstocks¹³². Biochar's impact on soil quality, nutrient availability, and soil carbon mineralization, however, varies, with studies reporting different outcomes^{133–135}. For instance, applying biochar at a rate of 30 t ha⁻¹ resulted in a 14% decrease in SOC content, whereas application rates of 60 t ha⁻¹ and 90 t ha⁻¹ led to SOC increases of 18.8% and 8.2%, respectively. The net change in SOC is therefore influenced by complex soil dynamics, necessitating a deep understanding of the target soil.

Soil fungi also present another avenue for increasing stable SOM. For instance, plants allocate up to 20% of their carbon to fungal partners via root exudation and detritus¹³⁶, effectively transforming labile organic compounds into more stable SOC. Sclerotia, the hardened masses of fungal hyphae, which typically range from 0.1 mm to 30 cm in diameter, can persist in soil for many years^{137,138}. Melanized endophytic fungi may also support this endeavour^{139–141} and have been shown to enhance polyaromatic SOC associated with microaggregates and slow decomposition of SOC, potentially increasing the recalcitrant pool of soil carbon^{139,140,142}.

Discussion S7. Figure data sources

Figure 1 | Major landmarks of the global carbon (C) cycle present opportunities for scalable intervention on atmospheric carbon accumulation. Global annual emissions from land use change and net anthropogenic emissions (and their year on year changes) were taken from the Global Carbon Budget report¹⁴³. Global annual heterotrophic CO₂ fixation is based on the midpoint figure derived from upper and lower bounds estimated in¹⁴⁴. Global annual heterotrophic respiration is the average of values reported in^{145–147}. The average global annual uptake from photosynthesis is based on values reported in^{148–151}, year on year change in photosynthesis is taken from¹⁵⁰. Global annual emissions from plant respiration is based on¹⁵². Global annual carbon removal from the combined CDR programs and the associated year on year change are taken from the 2nd Edition of The State of Carbon Dioxide Removal¹⁵³. Gross global annual ocean-atmosphere CO₂ absorption and emissions totals are taken from¹⁵⁴, the year on year change for ocean-atmosphere exchanges (absorption is taken from¹⁵⁵). The estimated magnitude of the oceanic biological carbon pump is presented as the midpoint of the accepted range presented in¹⁵⁶. Oceanic and

terrestrial (biotic) carbon pool sizes are based on values presented in ¹⁵⁷ whereas the terrestrial (pedologic) pool size is as reported in¹⁵⁸. Atmospheric carbon pool size was calculated (using the formula *Total atmospheric C pool size (GtC) = Atmospheric CO₂ concentration (ppm) × 2.12 GtC/ppm*). Atmospheric CO₂ concentration was based on the NOAA recorded atmospheric carbon levels data (taken from <https://climate.nasa.gov/vital-signs/carbon-dioxide/>), which was 425 ppm at the time of writing.

Figure 2 | Carbon sequestration strategies carry broad sequestration lifetimes and involve a diverse variety of biological systems. Data sources described above, see Discussion S2 and related tables.

Figure 3 | Dynamics in oceanic carbon cycling and overview of oceanic CCS approaches. Pools with known magnitudes are presented (using mean values where ranges have been previously reported) and processes with known rates are also given, data sourced from ^{159–164}.

Figure 5 | Biological systems drive compositional dynamics of the atmosphere via the balanced interplay of interlinked biogeochemical cycles. Global mean surface temperature (GMST) together with corresponding CO₂ levels over the past 485 Ma was adapted from Judd et al. 2024¹⁶⁵. Eccentricity data were collected using the NASA Goddard Institute for Space Studies ModelE AR5 simulation tool at <https://data.giss.nasa.gov/modelE/ar5plots/srorbpar.html>. Ice core data was obtained from Luthi et al. 2008¹⁶⁶ for CO₂, Loulergue et al. 2008¹⁶⁷ for CH₄, and Schilt et al. 2010¹⁶⁸ for N₂O.

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