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1 **The spectrum of slip behaviours of a granular fault**
2 **gouge analogue governed by rate and state friction.**

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6 **Key Points:**

- 7 • Slip modes in granular gouge are akin to natural fault slip.
8 • Glass beads are a suitable granular analogue for fault gouge and show rate-and-
9 state dependent friction.
10 • Enhanced creep and small scale events are signals for imminent failure and indi-
11 cate fault criticality.

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Abstract

The exact principles of earthquake recurrence and magnitude are currently unknown which is why earthquake hazard assessment relies on statistical models combined with numerical simulations. A component of seismic and aseismic slip is the frictional character of a fault. We shear fused glass beads with a narrow particle size distribution of 300-400 μ m at stresses of 5-20kPa and with low shear rates of less than 1mm/s. As a result, we show that characteristic slip events emerge, ranging from fast and large slip to small scale oscillating creep and stable sliding. In particular we observe small scale slip events that occur immediately before large scale slip events for a specific set of experiments. Similar to natural faults we find a separation of scales by several orders of magnitude for slow events and fast events. Enhanced creep and transient dilatational events pinpoint that the granular analogue is close to failure. From slide-hold-slide tests, we find that the rate-and-state properties are in the same range as estimates for natural faults and fault rocks. The fault shows velocity weakening characteristics with a reduction of frictional strength between 0.8 to 1.3 % per e-fold increase in sliding velocity. Furthermore, the slip modes that are observed in the normal shear experiments are in good agreement with analytical solutions. Our findings highlight the influence of micromechanical processes on macroscopic fault behaviour. The comprehensive dataset associated with this study can act as a benchmark for numerical simulations and alleviate the understanding of observations of natural faults.

Plain Language Summary

Earthquakes occur when two continental plates slide along each other. The motion is concentrated at the interface of the two plates which is called a fault. In many cases the fault is filled with granular material, called gouge, that supports the pressure between the plates. Therefore, the properties of this gouge determine how fast and how large an earthquake can be. It also has an influence on the time between earthquakes. In our study we examine a simplified version of a fault gouge in a simple small-scale model. Instead of rock material we use glass beads and measure how different conditions affect the motion of the model. We find that our model reproduces features of fault gouge because it shows similar behaviour. When there is no motion our model fault becomes stronger with a rate equal to fault gouge. Also, the type of strengthening is analogous to fault gouge. During slip, the glass beads become weaker as the slip velocity increases in a similar manner as natural faults. These results improve the understanding of computer simulations and natural observations.

1 Introduction

Seismically active faults pose a major threat to many communities world-wide. Therefore, it is vital to make appropriate predictions on the probability of large earthquakes and their associated effects, such as tsunamis and mass movements. Several factors contribute to the difficulties to estimate seismic hazard in the vicinity of such faults. Besides the vulnerability of structures and the societal impact, geological factors play an important role in seismic hazard assessment and the development of models that describe fault activity (Zöller & Hainzl, 2007). Current models for earthquake recurrence incorporate mathematical models of earthquake statistics (Gutenberg-Richter, Omori-Utsu-Aftershocks, Brownian-First-Passage-Time), numerical models of earthquakes and rupture processes (Rate-and-State-Friction), interseismic stress built-up and the interaction of multiple faults over a larger area via stress transfer (e.g. Brinkman et al., 2016; Ellsworth et al., 1999; Field et al., 2014; Hainzl et al., 2013; Hu & Bradley, 2018; Kawamura et al., 2012; Lapusta & Rice, 2003; Parsons, 2005; Zöller et al., 2011). These models inherently rely on the accurate description and characterization of fault properties and behaviour, as well as extensive catalogues of slip events. With this study we aim to characterize a

62 physical scale model of seismic activity to expand models of seismic hazard assessment
 63 with experimental data and also show the potential impact of various slip modes on seis-
 64 mic activity.

65 1.1 Fault Slip

66 Active faults are characterized by a wide range of slip behaviours ranging from aseis-
 67 mic creep to seismic stick-slip that may change spatially along the fault and temporally
 68 over the seismic cycle (e.g. Harris, 2017; Peng & Gomberg, 2010). The types of slip are
 69 defined by their characteristic timescale which ranges from milliseconds to a few years
 70 (Obara & Kato, 2016) and by their characteristic magnitude which is usually defined by
 71 seismic moment (Ide et al., 2007; Gomberg et al., 2016). Depending on their character-
 72 istics in time and seismic wave forms, the slip events are characterized as seismic (very
 73 low frequency earthquakes, tremors, normal earthquake) or geodetic (short-term and long-
 74 term slow slip events) events. They can occur simultaneously, i.e. within one seismic cy-
 75 cle, at the same locality or in different depth ranges of the same main fault (Bürgmann,
 76 2018). The physical origin of this range of slip modes is still not entirely clear, although
 77 several approaches for certain phenomena have been proposed (Daniels & Hayman, 2008;
 78 Ciamarra et al., 2010; Chen & Spiers, 2016; Dorostkar & Carmeliet, 2018).

79 A common methodology to model this wide range of slip behaviours is through a
 80 continuum based description that reproduces the kinematics and dynamics of fault ac-
 81 tivity. The rate-and-state framework provides the possibility to characterize fault behaviour,
 82 or in a general term 'fault rheology', by describing the connection of forces in the sys-
 83 tem (friction μ) and the external influences such as loading rate v_L and stiffness k (Brace
 84 & Byerlee, 1966; Dieterich, 1978; J. H. Dieterich, 1979a, 1979b; Scholz, 1998). In gen-
 85 eral, the rate-and-state framework is able to describe most observations that lead to fault
 86 (in-)stability and has been derived from experimental observations in the laboratory and
 87 a few field observations (Marone, 1998, and references therein). Stick-slip experiments
 88 using rock and rock analogues suggest that besides intrinsic material properties (e.g. fric-
 89 tion coefficient, slip/velocity weakening), extrinsic parameters like stiffness, normalized
 90 loading rate and effective normal stress are key controls of frictional stability (e.g. Lee-
 91 man et al., 2016; Heslot et al., 1994; Marone, 1998; Mair et al., 2002). Recent studies
 92 also highlight that several of the fault intrinsic parameters in the rate-and-state equa-
 93 tion are also dependent on extrinsic parameters and not constants as previously assumed
 94 (Van den Ende et al., 2018; Chen & Spiers, 2016).

95 1.2 Granular Fault Analogues

96 In this study we purely focus on the frictional characteristics of an analogue fault
 97 zone which is described with the rate-and-state framework (J. H. Dieterich, 1979a; J. Di-
 98 eterich, 2007). Our fault zone is composed of a granular fault core with relatively stiff
 99 outer boundaries and dominated by granular mechanics. Other processes that influence
 100 the slip modes along a fault zone, which are not realized in our setup, are variations in
 101 pore-fluid pressure, changes in material because of comminution, or mineral reactions.
 102 Not all slip modes are observed for all active zones which strongly suggests that there
 103 is a complex interaction between the processes acting on different scales in space and time.
 104 Knowledge of the complex interactions between the different slip modes is relevant for
 105 estimating the seismicity rates along plate boundaries and therefore for seismic hazard
 106 assessment. Other possible areas of application include soil mechanics and mass move-
 107 ments.

108 The advantage of using a granular analogue is the simplicity with which observa-
 109 tions can be made. The analogue modelling approach features lower stresses which sim-
 110 plifies the design and construction of the testing machine. This increases the available
 111 parameter space because it is relatively easy to change the system stiffness using springs.

112 For rock mechanical testing apparatuses the change in stiffness is limited to a smaller
 113 range that is either accessible through adding rubber blocks or by artificially changing
 114 the servo-hydraulic systems to mimic a different stiffness (Beeler et al., 1994) The re-
 115 sults from this study can be used to improve current numerical models of granular gouge
 116 but can also directly be applied in improved seismotectonic scale models (Rosenau et al.,
 117 2017; Blank & Morgan, 2019). The glass beads show a slip behaviour that naturally emerges
 118 from their frictional properties. This can be exploited for larger analogue models to model
 119 fault slip in a geometrically complex fault system. The range of available temporal and
 120 spatial scales, as well as the self-consistent scaling behaviour allow the application in many
 121 fields where rate-and-state friction is a dominant process such as landslides, glacial mo-
 122 tion, mass movements and lithospheric deformation (Jerolmack & Daniels, 2019). In com-
 123 parison to numerical simulations the use of an analogue model allows to inherently link
 124 the spatial and temporal scales without having to rely on parametrization and grid based
 125 methods. The analogue approach allows to model small scale processes, such as earth-
 126 quakes within a fault zone over many seismic cycles and over a much larger spatial scale
 127 within a shorter period of time than numerical simulations of similar complexity.

128 We here report characteristics of slip events in an analogue fault gouge consisting
 129 of spherical glass beads. In contrast to similar experiments of Frye and Marone (2002);
 130 Anthony and Marone (2005); Ferdowski et al. (2013); Jiang et al. (2016); Cui et al. (2016)
 131 we explore the low pressure (kPa instead of MPa) regime which is rich in slip behaviours
 132 and generates regular stick-slip with more complete stress drops similar to seismic cy-
 133 cles along major faults in a highly reproducible and accessible way. Several studies es-
 134 tablished the large diversity in slip modes in such experiments. Changes in stiffness and
 135 normal stresses lead to first order changes in frictional stress, such as transition from stick-
 136 slip to oscillation and stable sliding (Heslot et al., 1994). Nasuno et al. (1997) found lo-
 137 calized precursor phenomena in thin sheared glass beads that precede large slip events.
 138 Moreover, the use of a ring-shear tester instead of commonly used direct shear appara-
 139 tuses allows us to apply an in principle infinite amount of displacement and therefore
 140 a large number of events, which is a solid database for statistical analysis. Results from
 141 a similar apparatus by Cain et al. (2001) show that it is suitable to measure dilation-
 142 compaction cycles and show that the conditions in an annular shear cell lead to dilation
 143 during the loading phase and compaction during failure which is similar to the results
 144 obtained from rock mechanical tests in biaxial compression setups (Beeler & Tullis, 1997).

145 For the same material we vary the extrinsic parameters normal stress σ_N , loading
 146 velocity v_L , and stiffness k_L . In this parameter space, we monitor the occurrence of slip
 147 events and creep, as well as the transitions from one slip mode to another. We charac-
 148 terize the analogue fault gouge with commonly used tests to derive the rate-and-state
 149 parameters, such as slide-hold-slide tests (SHS). We compare the findings to first order
 150 observations from rock friction experiments and assess the suitability of granular ana-
 151 logue fault gouge for its use in combined analogue and numerical modelling.

152 2 Methodology

153 To simulate fault behaviour in various settings we use a granular analogue mod-
 154 elling approach. Previous studies examined granular media under natural pressure con-
 155 ditions, whereas we are using conditions realized by analogue models, being 3 to 4 or-
 156 ders of magnitude lower (Rosenau et al., 2017). This prevents comminution of the glass
 157 beads and ensures constant frictional properties over the experimental duration, which
 158 gives well reproducible results. In a first step the data is analysed with simple methods
 159 to quantify basic properties which makes it possible to easily compare the results with
 160 previous works. The terminology for certain points and characteristics of the data is found
 161 in appendix Appendix A and Tab. A1. The data analysis is done using a suite of Python
 162 scripts that pick events and do statistic calculations. All of which are going to be avail-

163 able as the open source software 'RST-Stick-Slip' from the GFZ git repository (Rudolf,
164 in prep) and are also included in the data publication (Rudolf et al., in prep).

165 2.1 Rate-and-State Friction

166 The relation between shear stress and normal stress for granular media and many
167 other interfaces is determined by a non-linear combination of mean stress, slip velocity,
168 stiffness and several non-dimensional parameters. This relationship is termed rate-and-
169 state dependent friction that macroscopically leads to alternating cycles of slip, creep
170 and locking, called stick-slip (Dieterich, 1978; J. H. Dieterich, 1979a; Ruina, 1983; Marone,
171 1998; Tullis & Weeks, 1986; Beeler et al., 1994). This effect is used to describe and ex-
172 plain the various slip behaviours that are associated to earthquakes, e.g. slip on faults
173 (Marone, 1998), earthquake nucleation (J. H. Dieterich, 1992) and slow slip events. In
174 our study we use the relationships and testing procedures defined in Beeler et al. (2001),
175 Marone and Saffer (2015) and J. Dieterich (2007) as well as adapted methodologies of
176 Corbi et al. (2013), Bhattacharya et al. (2015) and Bhattacharya et al. (2017) to esti-
177 mate the principal parameters for the rate and state equation. A short description of rate-
178 and-state friction and its application to our study is found in appendix A2.

179 To test which of the state evolution laws best describe our experimental data we
180 take a semi-quantitative approach which considers certain observations, such as the evo-
181 lution of stress during a hold phase, or the behaviour in unstressed SHS-tests. These re-
182 sults are then compared to other experimental findings. In some cases a close quanti-
183 tative comparison is possible, while in others either the experimental setups are too dif-
184 ferent for a direct comparison, or an easily comparable quantity could not be found. E.g.
185 the general dilatational behaviour in the reloading phase just after a hold period is simi-
186 lar for our experiments compared to the results by Beeler and Tullis (1997) which can
187 then be qualitatively interpreted in the context of state evolution (Bhattacharya et al.,
188 2017) (see section 4.2). Strong stick-slip effects, probably due to insufficiently high ma-
189 chine stiffness, prevented a direct fit of Eq. A1 to the data to estimate rate-and-state
190 parameters from classical velocity stepping.

191 2.2 Experimental Setup

192 For the experiments we use a ring shear tester of type 'RST-01.pc' (Schulze, 1994;
193 ASTM, 2016) which allows to shear a granular sample in an annular shear cell. The ma-
194 chine and methodology has been verified and calibrated using a standard bulk material
195 (CRM-116 limestone powder) and is extensively used for characterizing granular mate-
196 rials in engineering and analogue modelling (e.g. Ritter et al., 2016a; Klinkmüller et al.,
197 2016; Schulze, 1994).

198 The granular material is confined in a ring shaped shear cell and sheared against
199 a lamellae-casted lid which also imposes the normal load (Fig. 1a+b). The normal load
200 is adjusted using a motorized weight attached to a lever that pulls the lid from below.
201 This ensures a constant normal load on the sample. Two bars attached to force trans-
202 ducers hold the lid in place and measure the shear forces acting on the lid.

203 The applied and resulting forces (normal and shear), driving velocity and vertical
204 lid displacement are measured as individual channels at the analogue output of the ma-
205 chine. The main set of experiments were measured using a Peripheral Component In-
206 terconnect (PCI) based analogue-to-digital converter card (ADC) at a frequency of 12.5
207 kHz each (BMCM - PCI Base 50, controlled with BMCM Nextview[®] software). The mea-
208 sured values are averaged over 20 samples for noise reduction resulting in a final output
209 frequency of 625 Hz. Another set of experiments, mainly the slide-hold-slide (SHS) tests,
210 were measured with a real-time embedded controller (NI - CompactRIO) at 50 kHz per
211 channel using an ADC module (C Series Universal Analogue Input Module, NI-9219, con-

212 trolled by custom in-house software). This change was due to the end-of-life of the op-
 213 erating system during the course of this study which lead to hardware incompatibilities
 214 with the PCI-based approach. Similar to the other measurements, this data is averaged
 215 down to a frequency of 1 kHz. Based on the setup geometry, shear and normal forces
 216 are converted into shear and normal stresses according to ASTM (2016) and lid displace-
 217 ment into volumetric change (dilation/compaction). Shear forces are converted to shear
 218 stress using the moment of the crossbar $M_d = r_s \cdot F_s$, the median radius r_m and the
 219 cross-sectional area of the lamellae a_d (Eq. 1).

$$\tau = \frac{M_d}{r_m \cdot a_d} \quad (1)$$

220 The median radius r_m is also the reference position at which the loading rate v_L
 221 is defined because it divides the cell into two regions of equal volume. While the nor-
 222 mal stresses are also continuously measured, we assume a constant normal stress that
 223 is equal to the value set at the start of the measurement. Due to internal correction fac-
 224 tors which are not disclosed by the manufacturer there is always a slight discrepancy be-
 225 tween set normal stress and measured normal stress. Partially, this discrepancy stems
 226 from the angle of the tie rods and the lid which exerts additional normal stress onto the
 227 sample (pers. comm. D. Schulze).

228 As granular material we use 300-400 μm sized fused soda-lime glass micro-beads
 229 supplied from Kuhmichel Abrasiv GmbH (Fig. 1c). They are characterized by a rela-
 230 tively low dynamic friction coefficient ($\mu = 0.47$) and no measurable cohesion ($C =$
 231 $1 \pm 12 \text{ Pa}$) as well as a strain hardening-weakening behaviour associated with dilation-
 232 compaction (Lohrmann et al., 2003; Klinkmüller et al., 2016; Ritter et al., 2016a). Glass
 233 beads are frequently used as a rock and gouge analogue material and generate stick-slip
 234 under laboratory conditions (e.g. Mair et al., 2002). Because they are non-cohesive we
 235 can approximate the instantaneous frictional resistance of the fault zone μ as the ratio
 236 of shear stress τ to the applied normal stress σ_N :

$$\mu = \frac{\tau}{\sigma_N} \quad (2)$$

237 Before an experiment is started, the sieved samples are presheared by 10 mm at
 238 a loading velocity of $0.5 \frac{\text{mm}}{\text{s}}$ which ensures a fully developed shear zone without major
 239 post failure weakening (derived from Ritter et al., 2016a, 2016b). Tab. 1 lists the exper-
 240 imental parameters for the various tests performed for this study. For all tests (main and
 241 SHS) we use 4 different normal stresses of 5, 10, 15, and 20 kPa. The major difference
 242 between the tests is the stiffness k_M and loading rate v_L .

243 The main tests are conducted with logarithmically spaced velocity v_L from 0.02
 244 $\frac{\text{mm}}{\text{s}}$ to $0.0008 \frac{\text{mm}}{\text{s}}$. The duration of each run is the inverse of the respective loading ve-
 245 locity leading to equal displacement and a similar amount of slip events. Each individ-
 246 ual test is carried out at constant normal load.

247 To limit the influence of stick-slip we perform the stressed SHS tests with maxi-
 248 mum machine stiffness and at higher shear rates. To also estimate the rate effect on heal-
 249 ing rate we vary the loading rate v_L from 0.05 to $0.52 \frac{\text{mm}}{\text{s}}$. The hold times were increased
 250 logarithmically $t_{hold} = \{1 \dots 1000\} \text{ s}$ (in accordance with Eq. A7) and a constant load
 251 point displacement of 5 mm between the hold phases was applied. Additionally, we did
 252 one series of unstressed SHS tests with the same set of parameters as for the stressed SHS
 253 test but only with a single load point velocity of $v_L = 0.16 \frac{\text{mm}}{\text{s}}$. For unstressed SHS
 254 tests we reduce the normal stress to zero during hold to evaluate whether the change in
 255 state θ is purely time dependent (Aging law) or shows a slip dependent behaviour (Slip
 256 law). Due to the time needed for the machine to unload and reload the normal stress,

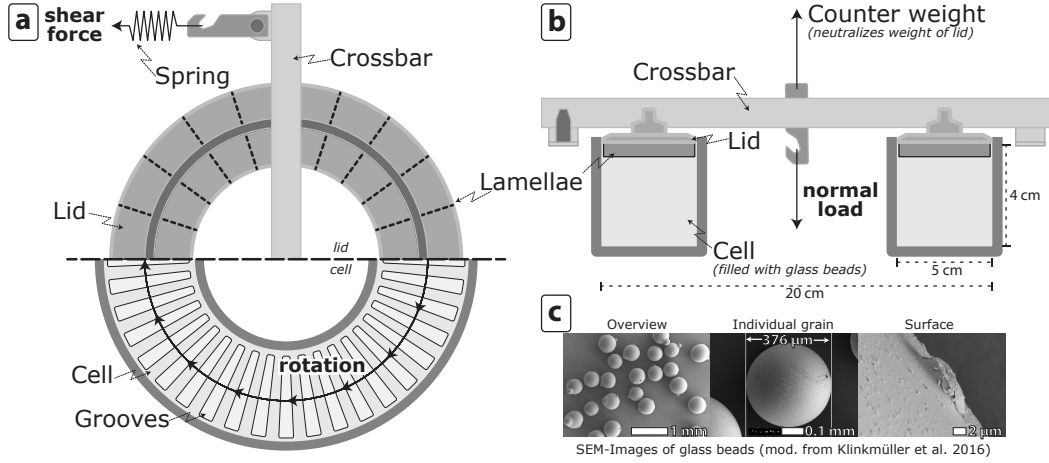


Figure 1. Schematic drawing of the modified ring shear tester. The system is loaded at loading velocities of $0.02 \frac{mm}{s}$ to $0.0008 \frac{mm}{s}$ by rotating the cell. The cell has grooves for a high friction interface which is mirrored by lamellae attached to the lid. A moveable weight pulls the lid from below by a motor driven lever for applying normal load. Force transducers behind the springs measure shear force. a) Top view the above part showing the lid and the bottom part showing the cell and its internal structure. b) Cross section through the whole setup. c) Scanning electron microscopy images of the glass beads showing the average particle size and the surface structures (modified from Klinkmüller et al., 2016).

257 which was a few seconds per kPa, we only used hold intervals of $t_{hold} = \{100...1000\} s$.
 258 All SHS-cycles were repeated three times to be able to estimate the variance of the mea-
 259 surements.

260 **2.2.1 Adjustment to other setups**

261 To compare our experimental conditions with other setups in terms of stress condi-
 262 tions and stiffness we utilize the normalized stiffness k_N and normalized loading rate
 263 v_N . Both have an influence on the material behaviour because of rate-and-state friction
 264 and are dependent on the setup. For our setup the normalized stiffness is calculated from
 265 the machine stiffness k_M , which is modified using springs, the geometrical factor L_M con-
 266 verting shear force to shear strain and the applied normal stress σ_N .

$$k_N = \frac{k_M \cdot L_M}{\sigma_N} \quad (3)$$

267 From this the normalized loading rate v_N is derived with the loading rate v_L (Eq.
 268 4). It can be interpreted as a non-dimensional stressing rate that describes the increase
 269 in stress counteracted by friction over time.

$$v_N = k_N \cdot v_L \quad (4)$$

270 For our setup the values for k_N are in the range of 10^{-2} to $10^2 mm^{-1}$ and for v_N
 271 varies between 10^{-5} and $10^{-1} s^{-1}$. This is in the same regime as Nasuno et al. (1997)
 272 and at least three orders of magnitude higher than the values achieved by Beeler et al.
 273 (1994).

Table 1. Experiment overview for this study

Type	Stiffness $[k] = \frac{N}{mm}$	Normal stresses $[\sigma_N] = Pa$	Load point velocities $[v_L] = \frac{mm}{s}$
Main tests (RST)	{3.3, 19.6, 82.6, 1354.0}	{5000, 10000, 15000, 20000}	{0.0008, ..., 0.02}
stressed SHS tests	1354.0	{5000, 10000, 15000, 20000}	{0.05, ..., 0.52}
unstressed SHS tests	1354.0	{5000, 10000, 15000, 20000}	0.16

3 Results

We here describe the slip modes qualitatively (Fig. 2 + 3) and quantitatively using the asymmetry of the event cycles (Fig. A2). Then we determine the constitutive parameters for our setup and analogue fault gouge which determine the stick-slip characteristics and slip behaviour by slide-hold-slide tests. To compare the data across the individual setups we use the normalized loading rate v_N (Eq. 4) as a key parameter. This parameter contains the joint influence of normal stress and loading rate and makes it possible to plot all experiments together without major overlap. Nevertheless, in all cases there is a distinct influence of both parameters for the individual datasets so that in the following sections the results as a function of normal stress and stiffness are presented as well. We use a similar colour and marker code in most plots that show results from the experiments. Normal stress, in some cases loading velocity, is indicated by colour while the setup stiffness k_M is indicated by markers. All errors in plots or in numeric values are given as twice the standard deviation (2σ) of the respective quantity. A rigorous error propagation is done during data analysis using the Python module 'uncertainties' (Lebigot, 2021).

3.1 Slip Mode

The slip mode is qualitatively defined by the evolution of stress during an experimental run (Fig. 2). Low stiffness leads to typical sawtooth shaped curves with very sharp acceleration immediately before failure (Fig. 2a₁). Increasing the stiffness increases the amount of pre-slip and slows the acceleration before failure. This is expressed as slightly smoother sawtooth curves but the duration of a slip event is still much shorter than the reloading phase (Fig. 2b₁). For Spring C we find oscillations of weakly irregular shape (Fig. 2c₁). On average the increasing edge of an oscillation is a bit longer but only by a factor of 1.5 to 2 and not several orders of magnitude as for the softer springs. Another slip mode is observed for the highest stiffness which shows stick-slip cycles with a plateau of high stresses before failure. If the sample is at these high stresses we observe small and slower slip events that occur very close to failure (Fig. 2d₁). Figure 3 shows an overview of all qualitatively determined slip modes in all experiments. Furthermore, it indicates the full phase space in stiffnesses k_M and loading velocities v_L that was surveyed in this study.

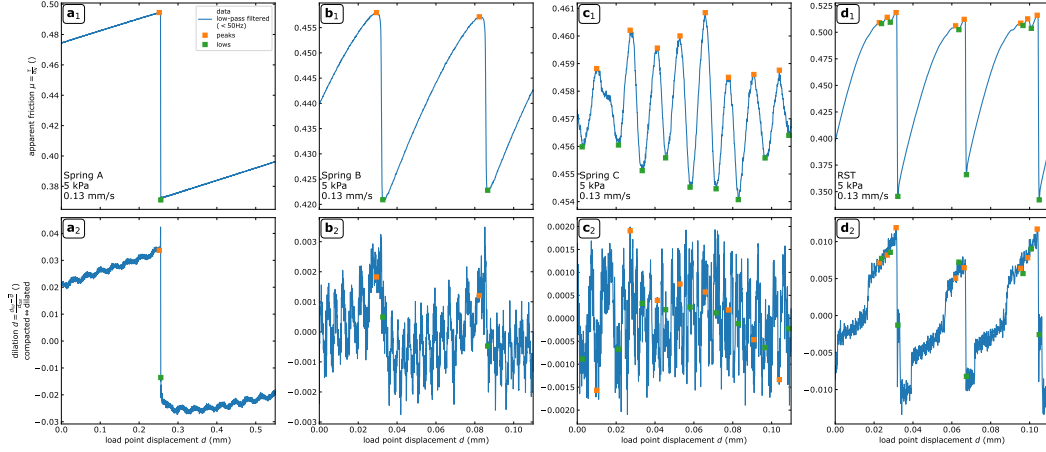


Figure 2. Exemplary stress and dilation curves during a typical experimental run. For all three experiments the normal stress and loading velocity are the same ($\sigma_N = 5 \text{ kPa}$, $v_L = 0.13 \frac{\text{mm}}{\text{s}}$) only the machine stiffness k_M increases from top to bottom. a) Spring A - $k_M = 3.3 \frac{\text{kN}}{\text{mm}}$: Regular sawtooth shaped stick-slip curve with a linear loading phase resulting from low stiffness. Due to the extremely high recurrence time this plot has been scaled down by a factor of 5 to be able to see a stick-slip event. b) Spring B - $k_M = 19.6 \frac{\text{kN}}{\text{mm}}$: Less sharp stick slip curves with a clear acceleration phase after peak strength. c) Spring C - $k_M = 82.6 \frac{\text{kN}}{\text{mm}}$: Oscillations due to intermediate stiffness. d) RST - $k_M = 1354.0 \frac{\text{kN}}{\text{mm}}$: Higher stiffness leads to non-linear loading behaviour and minor slip events just before major slip, while the loading velocity remains similar.

305 For all experiments we observe a characteristic succession of dilation and compaction.
 306 For perfect stick-slip, slip events lead to strong compaction of $\Delta d \approx 0.07$ grain diam-
 307 eters d_{GB} which then slowly dilates during the interevent period (Fig. 2a₂). In the first
 308 moments of failure for low stiffness experiments we often observe an initially dilating mo-
 309 tion in the first few milliseconds. Experiments with stick-slip also show oscillations and
 310 characteristic patterns on a variable scale while experiments in the oscillating regime show
 311 a mostly random pattern of $\pm 0.001 d_{GB}$ (Fig. 2c₂). For low stiffness experiments this os-
 312 cillation is on a scale of $\pm 0.003 d_{GB}$ with a period of 2.5 s. Increasing stiffness leads to
 313 a reduction of the oscillation period to values of 0.2 s for Spring B and < 0.1 s for RST.
 314 With increasing amounts of creep we find additional oscillations in the dilation signal
 315 which gradually change their period closer to failure (Fig. 2b₂). For bimodal experiments,
 316 we observe a pattern of abrupt dilation events during the interevent period which is sim-
 317 ilar for consecutive interevent phases (Fig. 2d₂). The gradual change in oscillation pe-
 318 riod is also present for these experiments but on a smaller scale than for lower stiffness.

319 Quantitatively we describe the slip mode through the average asymmetry r_a of the
 320 stick-slip curves and its distribution. Asymmetry r_a is defined as the ratio of slip dura-
 321 tion t_s versus the reloading time t_r (Fig. A2). We find that at generally low normalized
 322 loading rates ($v_N < 10^0 \text{ s}^{-1}$, low stiffness and low normal stress, Fig. 3a+b, Fig. A2a)
 323 the asymmetry is very high and has a low variability, although the dataset is relatively
 324 sparse in that region due to very long reloading phases ($t_{rel} > 10^2 \text{ s}$, e.g. Fig. A2a₃).
 325 Pure stick-slip, dominant at low setup stiffnesses, has a very high asymmetry because
 326 the reloading phase is very long compared to the duration of an event (e.g. Fig. A2a).
 327 In general both experiments with low stiffness (Spring A and B), show relatively regu-
 328 lar and well defined stick-slip events. This is indicated by relatively flat increases in shear
 329 stress and abrupt decays with strong to modest acceleration before failure (Fig. 2a+b).

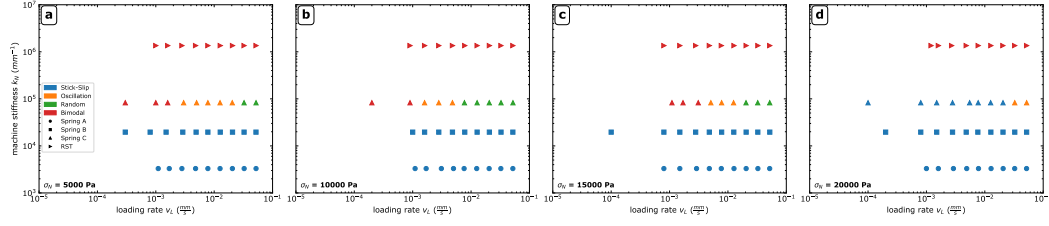


Figure 3. Qualitatively determined slip modes for all experiments in the full k - v -space. All experiments for RST show a bimodal slip distribution and all experiments for Spring A and B show well defined stick-slip cycles. The intermediate stiffness for Spring C leads to variable slip modes depending on loading rate v_L and normal stress σ_N . A transition from bimodal via oscillations to random is found for increasing loading rate. At higher normal stresses the bimodal slip mode is replaced by stick-slip and is also present at higher loading rates. Each column displays the result for a different normal stress. The legend applies to all subplots.

330 Experiments with Spring C show three different slip modes depending on normal
 331 stress and loading rate. With increasing loading rate we observe an evolution from bi-
 332 modal over oscillation to random. This evolution is clearest for low normal stresses and
 333 less apparent than for high normal stress (Fig. 3). At normalized loading rates $v_N <$
 334 10^0 the slip mode is bimodal with oscillating events preceding asymmetric events (Fig.
 335 3a-c and e.g. Fig. A2c₁). In the interval $v_N = \{10^{-0.5} \dots 10^{0.5}\}$ we find low asymme-
 336 try with low variability which approaches $r_a = 1$, which is the expression of oscillat-
 337 ing events becoming more and more symmetrical. The asymmetry decreases until we find
 338 oscillating slip modes at normalized loading rates between 1 and 10 s^{-1} . Oscillations are
 339 characterized by symmetrical increases and decreases of shear stress with an almost si-
 340 nusoidal character (Fig. 2c). In terms of asymmetry this leads to an average ratio $r_a \approx$
 341 1 with a small variance (Fig. A2c₁₋₃). At higher normalized loading rate the system be-
 342 comes random and the asymmetry shows a large variance with a mean asymmetry of $r_a \approx$
 343 1. Slip under these conditions tends to be chaotic and does not show any characteris-
 344 tic features.

345 At highest stiffness $k_M = 1354.0 \frac{kN}{mm}$ (RST) we see a different evolution of slip
 346 mode with a complex sets that are influenced by normal stress and loading rate. For lower
 347 rates $v_N < 10^2$ the events split into two different distributions one with high asymme-
 348 try and one with lower asymmetry (Fig. A2d₂₋₄). As shown in Fig. 2d, slow and small
 349 events alternate with larger events in a characteristic sequence. The duality of slip modes
 350 leads to a bimodal distribution of asymmetry with a distinct separation of small/slow
 351 and large/fast events. Normal stress has a strong influence on how well defined the sep-
 352 aration between these two is. High normal stress leads to a clear separation which is at
 353 least one order of magnitude. At higher rates continuous distribution is observed while
 354 retaining a mean that is larger than 1, which still indicates defined stick-slip cycles rather
 355 than randomness.

356 In general, we find that above a certain stiffness the slip mode switches from simple
 357 stick slip (Spring A+B) to a more complex pattern of slip modes. Furthermore, high
 358 loading rates suppress the evolution of well defined stick slip cycles and lead to oscilla-
 359 tions and random slip modes. Normal stress defines the behaviour for low loading rates,
 360 which is most apparent at higher stiffness (Spring C + RST). At low normal stress, the
 361 slip mode is mainly bimodal which is due to higher amounts of creep. Increased normal
 362 stress suppresses creeping mechanisms and forces a change from bimodal slip mode, with
 363 slow events to pure stick slip. Additionally, this leads to a shift in mode space so that
 364 oscillations occur at higher loading rates (Fig. 3a-c vs. d).

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3.1.1 Stiffness

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The setup has two types of stiffness, one with and one without sheared material. The latter is straightforward to measure by fixing the lid to the shear cell and measuring the force increase while moving the shear cell. The basic stiffness of the apparatus $k_M = 1354.0 \frac{kN}{mm}$ (RST) is mainly influenced by the stiffness of the load cells that measure shear force which acts in series with the aluminium of the lid, tie rods and cross-bar. Adding springs in between the load cells and tie rods lowers this stiffness to the values reported in Tab. 1 (Spring A - $k_M = 3.3 \frac{kN}{mm}$, Spring B - $k_M = 19.6 \frac{kN}{mm}$ and Spring C - $k_M = 82.6 \frac{kN}{mm}$).

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For the types of tests reported in this study another type of stiffness is of relevance, the reloading stiffness k_L which is a combination of machine stiffness k_M and material stiffness. This property is calculated from the linear part of the reloading phase between slip events. We find that for experiments with low machine stiffness (Spring A+B) the normalized reloading stiffness k_R is roughly one order of magnitude smaller than the normalized machine stiffness (Fig. 4). For Spring B there are a few outliers and a weak increase in k_R is observed due to the influence of normal stress for the lowest normal stress $\sigma_N = 5kPa$. Spring C shows different results depending on the type of events considered. If only dynamic events with a slip rate above a critical threshold (comp. section 3.1) are considered (Fig. 4a) the reloading stiffness k_L is one order of magnitude smaller than the machine stiffness only for high normal stresses ($\sigma_N = 20kPa$), for lower normal stresses ($\sigma_N < 15kPa$) the difference is reduced to only half an order of magnitude. Considering all events (Fig. 4b) we see that especially for these lower normal stresses there is an influence of loading rate, that is apparent from the large spread in values due to variable slip modes (Sec. 3.1). The strongest difference is measurable for the highest machine stiffness (RST) where k_R is roughly 1.5 orders of magnitude smaller, with only a minor increase in variability for all events.

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3.2 Event Magnitudes

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Comparing the frictional stress drops $\Delta\mu$ for all experiments we find three different groups of stress drop highlighted in Fig. 4c. These are distinguished by their magnitude and variability of stress drop, as well as the evolution with increasing normalized loading rate.

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The first group occurs at low to medium normalized loading rate and at high stress drops (red area in Fig. 4c). The stress drop shows an exponential decrease with a similar slope for most experiments in this region. It consists mainly of fast slip events and experiments at low to intermediate stiffness (Spring A, B and C). A minor outlier is the Spring A-5 kPa experiment which has slightly higher stress drops but the same slope.

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The second group consists exclusively of fast events at the highest stiffness (blue area in Fig. 4c). They all plot at high normalized loading rate and show the highest stress drops which is roughly one order of magnitude higher than for the previous group at similar normalized loading rate. The evolution of stress drop shows a different slope, but is also decreasing with increasing normalized loading rate.

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The third group are slip events with a low stress drop $\Delta\mu < 10^{-2}$ and a large variability in stress drop that may span one or two orders of magnitude (green area in Fig. 4c). This group is dominated by slow events. In general, the stress drop is decreasing for increasing normalized loading rate, but the slope is not constant.

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3.3 Slip Velocities

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From the above observations we find that there is a characteristic difference between certain events under certain conditions leading to a bimodal distribution of asymmetry.

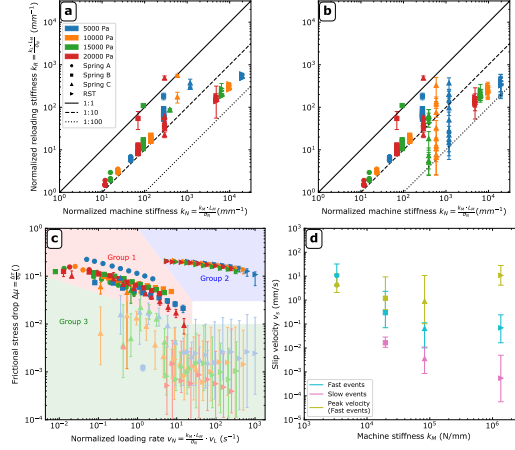


Figure 4. a+b) Measured reloading stiffness k_L in comparison with machine stiffness k_M for reloading phases of a) fast events and b) all events. The reloading stiffness k_L is one order of magnitude smaller than the machine stiffness (dashed line). c) Frictional stress drop $\Delta\mu$ distributions for all experiments. Each point represents the median and the error bars enclose 95% of all values. Solid colours highlight events which are found to be 'fast' events, 'slow events' are shown in lighter colour. The coloured areas define the individual groups that were identified. d) Average slip velocity during an event for fast events and slow events. The experiments with lowest stiffness (Spring A) show only fast events and the slow event data point for Spring B is based on a single experiment.

413 This difference is highlighted using the average slip velocity \bar{v}_s during an event as an in-
 414 dicator (Fig. 4d). The fastest slip events with $v_s \approx 10 \frac{mm}{s}$ are observed for the low-
 415 est stiffness (Spring A, $k_M = 3.3 \frac{kN}{mm}$) which is mainly due to the much larger slip dur-
 416 ing an event. With increasing stiffness the fast slip velocity is decreasing to a level of $10^{-0.5}$
 417 to $10^{-1} \frac{mm}{s}$. When slow events are present, which is not the case for all experiments at
 418 intermediate stiffness (Spring B+C), they are generally one to two orders of magnitude
 419 slower than the fast slip events ($v_s \approx \{10^{-2} \dots 10^{-3}\} \frac{mm}{s}$). The difference between fast
 420 and slow events increases towards the highest stiffness experiments, which also has the
 421 highest variability for slow slip velocities. Additionally, the peak slip velocity, that is the
 422 highest instantaneous slip velocity during a slip event, increases with increasing stiffness.
 423 The peak slip velocities are generally higher or in the same range as the mean slip ve-
 424 locity for fast events. Towards higher stiffness, the peak slip velocity is 2 orders of mag-
 425 nitude higher than the average slip velocity during a fast event. Typical peak slip ve-
 426 locities are in the range of $v_S = 1 \dots 10 \frac{mm}{s}$.

427 These slow events are characterized by low stress drop, low stress drop rate and
 428 a characteristic occurrence late in the cycle at generally high mean stress (Fig. 5). The
 429 relative amount of slow events decreases with increasing normal stress. For low normal
 430 stress more than 40% of the total events are found to be slow events, whereas for higher
 431 normal stresses it is 5 - 10%. Additionally, there is a variation in occurrence with load-
 432 ing velocity. At high loading velocity only very few slow events are detected, while at
 433 low loading velocity multiple slow events of increasing size can occur before one main
 434 event.

435 We find that in the series with Spring C ($k_M = 82.6 \frac{kN}{mm}$) at low normal stress and
 436 low loading rate these events show a nearly oscillating pattern of multiple cycles that is
 437 occasionally perturbed by a fast slip event. This is also highlighted in the asymmetry

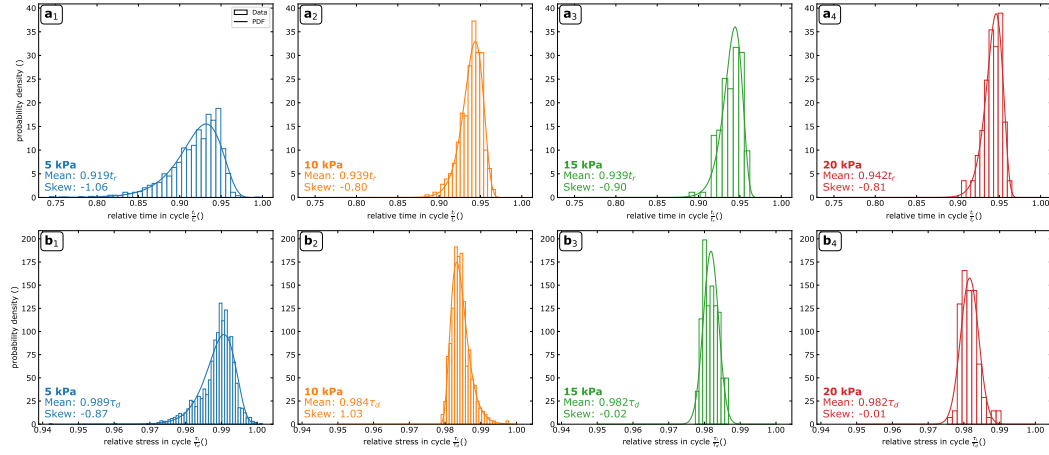


Figure 5. Timing and stress level of slow events during experiments of the highest stiffness ($RST, k_M = 1354 \frac{kN}{mm}$). a) Relative temporal occurrence of slow slip events. The probability increases towards failure with a maximum of $0.95t_r$ and very few events before $0.90t_r$ and after $0.96t_r$. b) Stress level where the slip events occur. The events are almost normal distributed with maxima between $0.98t_r$ and $0.99t_r$.

438 where these experiments have a bimodal distribution with one mode at high asymme-
 439 try (fast events) and one mode at low asymmetry (slow events). For the highest stiff-
 440 ness ($RST, k_M = 1354.0 \frac{kN}{mm}$) the slip rates are similar to Spring B for the fast events,
 441 but the slow events are slightly slower and show a higher asymmetry. There are fewer
 442 slow events and of smaller magnitude, with an average stress drop that is only 2.6% of
 443 the corresponding main event.

444 The occurrence of slow events shows a specific temporal pattern for the highest stiff-
 445 ness. The temporal distribution of slip events show a log-normal distribution skewed to-
 446 wards the end of the cycle and they do not occur in the first half of a cycle. The prob-
 447 ability of occurrence increases from $0.7t_r$ onwards with a mean of $0.92t_r$ to $0.94t_r$ and
 448 peaks at $\approx 0.95t_r$ (Fig. 5a). Then the probability drops abruptly to zero and for all ex-
 449 periments almost no precursor has been detected in the last moments of a cycle. Higher
 450 normal stresses shift the onset of occurrence closer to failure with a smaller variability
 451 but still with no events immediately before failure. The stress level at which the slow
 452 events occur is generally very close to the stress level of the main event (Fig. 5). The
 453 curves show a log normal distribution (skew ≈ 0) at higher stress levels which changes to a nor-
 454 mal distribution (skew ≈ 0) at higher stress levels which changes to a normal distribution (skew ≈ 0)
 455 occur around $0.98\tau_d$, and for $\sigma_N = 5$ kPa at higher levels of $0.99\tau_d$. Again we see a strong
 456 increase in occurrence up to a certain level of stress and an absence of events at stress
 457 levels very close to failure strength ($>0.99\tau_d$).

458 3.4 Interevent times

459 In experiments where a unimodal distribution of asymmetry is found, it is straight-
 460 forward to define the interevent time as the time between the individual events. But for
 461 bimodal distributions it is more complex. Therefore we use the term 'recurrence time
 462 t_{rec} ' for the time between any events (denoted by i) and the term 'reloading time t_{rel} '
 463 for the time between the fast events (denoted by d). This results in the following def-
 464 initions:

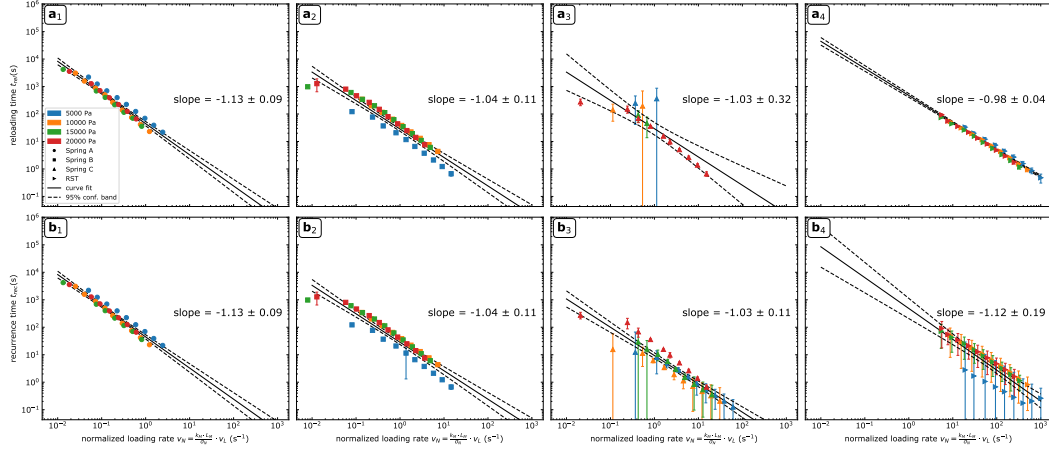


Figure 6. a) Reloading and b) recurrence times in comparison with normalized loading rate. The exponent is significantly different from $n = -1$ for the lowest stiffness (a_1 and a_2) which means that the recurrence decreases stronger than expected by the increase in normalized loading rate. The other stiffnesses show exponents that are only slightly smaller than $n = -1$ with larger errors.

$$t_{rec} = t_i - t_{i-1} \quad (5)$$

$$t_{rel} = t_d - t_{d-1} \quad (6)$$

465 In general the interevent times decrease with increasing normalized loading rate
 466 in an exponential fashion. The interevent times t_{rec} and t_{rel} are essentially the same for
 467 low stiffness setups (Spring A+B) because there is only one experiment in these series
 468 with slow events. For Spring C there are only few fast events which leads to a large error
 469 for the reloading times (Fig. 6a₃). Furthermore, there is a strong influence of slow
 470 event on the variance of recurrence times for the highest stiffness (RST, Fig. 6b₄). The
 471 power law exponent is slightly lower than $n = -1$ which means that there is a stronger
 472 decrease in reloading or recurrence time than what would be expected if there would be
 473 a direct correlation. Only the evolution of reloading time, that is for fast events, for the
 474 highest stiffness shows a power law exponent of $n = -0.98 \pm 0.04$ which indicates that
 475 the occurrence of fast events is directly proportional to the normalized loading rate.

476 3.5 Rate-and-State Parameters

477 The healing rate b is determined from the change in peak stress after increasingly
 478 larger hold times (Eq. A6). We find that all stressed SHS-tests show a positive healing
 479 rate (Fig. 7a). The mean healing rate from all fits is $b = 0.0057 \pm 0.0005$ which indi-
 480 cates time-dependent strengthening of the granular fault over time. There is no appar-
 481 ent correlation of healing rate b with loading velocity v_L (Fig. 7c). However, we observe
 482 a higher healing rate for a low normal stress of $\sigma_N = 5kPa$. Statistically it is not sig-
 483 nificantly different in the 95% confidence band in comparison with the other series due
 484 to a relatively high error for all fits at this normal stress ($\overline{b_{5kPa}} = 0.007 \pm 0.003$, Fig.
 485 7a₁). But the individual b values all plot outside the 95% interval of the mean fit of all
 486 data sets combined (black dotted line in Fig. 7d).

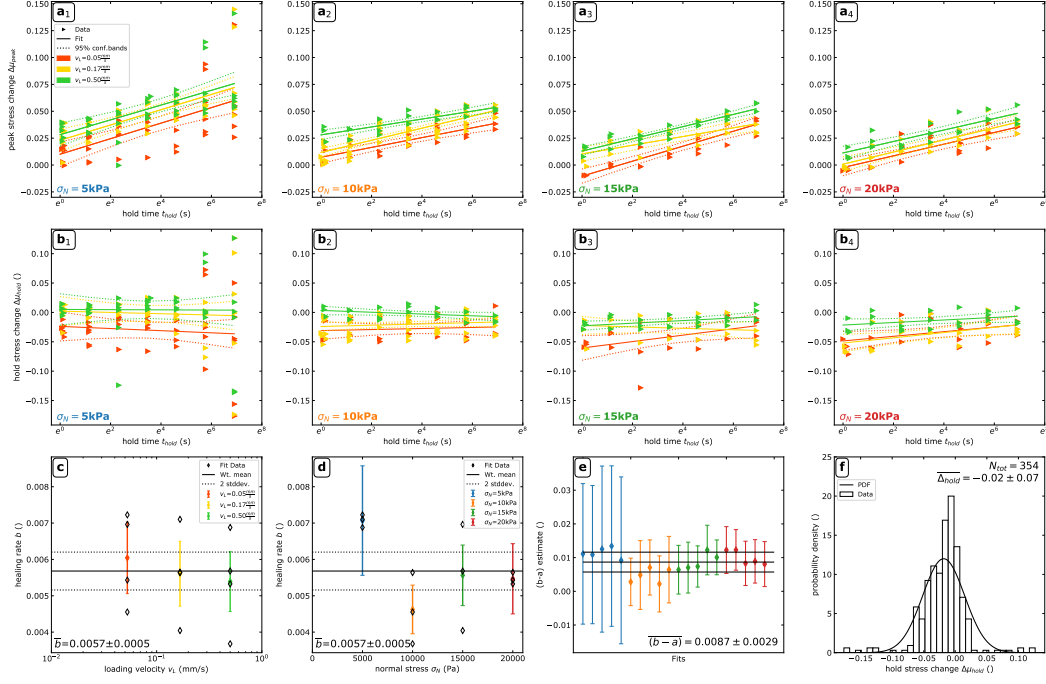


Figure 7. Overview of all slide-hold-slide related tests and quantities. a₁₋₄) Change in peak stress $\Delta\mu_{peak}$ after a hold interval t_{hold} compared to the average pre-hold level during sliding. The slope of the log-linear fit is the healing rate b which is positive for all experiments. The legend in a₁ applies to plots a₁₋₄ and b₁₋₄, the errors given are derived from the covariance of fit ($2\sigma = 2\sqrt{s^2}$). b₁₋₄) Change in hold stress $\Delta\mu_{hold}$ during a hold interval due to creep. c) Synthesis of all fits for healing rate b from a₁₋₄ with respect to loading velocity. The fitting errors on the 'Fit Data' points have been hidden for better visualization but are included in the error of the mean through error propagation with weighted averages. d) Synthesis of all fits for healing rate b from a₁₋₄ with respect to normal stress showing anomalously high values for $\sigma_N = 5$ kPa. The errors in this plot are displayed in the same way as for c). e) Estimation of $(b - a)$ from subsets of the experiments in a₁₋₄ sampled according to Eq. A7. The legend is the same as in d). f) Histogram of all hold stress changes $\Delta\mu_{hold}$ from b₁₋₄ showing a normally distributed change which is not significantly different from $\Delta\mu_{hold} = 0$ due to the high error.

487 The direct effect a is derived from two approaches. The first uses the offset in y-
 488 axis intersect of the peak stress change $\Delta\mu_{peak}$ (Eq. A8). This effect is clearly visible
 489 in Fig. 7a where the average peak stress change increases consistently for increasing load-
 490 ing rates while the slope stays constant. From the average increase in peak stress change
 491 with increasing loading velocity we compute a direct effect $a = -0.0074 \pm 0.0031$. As
 492 a result we calculate a first $(b - a) = 0.0131 \pm 0.0031$ from this observation only. The
 493 second approach exploits the selection of loading velocities with respect to the hold times
 494 so that Eq. A7 is fulfilled (after Beeler et al., 2001). The average $(b - a) = 0.0087 \pm$
 495 0.0029 fitting all possible combinations from all experiments (Fig. 7e). Using b from the
 496 previous estimate we arrive at a direct effect $a = -0.0030 \pm 0.0030$ which is less than
 497 the previous estimate.

498 Another important observation for rate-and-state friction is the change of stress
 499 during the hold phase $\Delta\mu_{hold}$ (Fig. 7b_{1-4+f}). The dataset is very noisy for this obser-
 500 vation. We observe a weak correlation of hold stress change over time with increasing
 501 normal stress. At low normal stress ($\sigma_N = 5kPa$, Fig. 7b₁) the data set shows a nega-
 502 tive slope which becomes smaller at $\sigma_N = 10kPa$ (Fig. 7b₂) and changes to a posi-
 503 tive slope for $\sigma_N \geq 15kPa$ (Fig. 7b₃₊₄). However, the estimated errors for these fits
 504 are quite large and while on average the hold stress change is negative $\Delta\mu_{hold} = -0.02 \pm$
 505 0.07 it is not significantly different from zero (Fig. 7f).

506 3.5.1 Additional Observations from Main Experiments

507 We observe an increase in peak strength with increasing reloading time for the main
 508 experiment series. Plotting the reloading time t_{rel} against peak frictional strength at fail-
 509 ure τ_p a log-linear increase can be observed (Fig. 8a). The observed slope ranges from
 510 $\beta = 0.0083$ to $\beta = 0.0130$ and indicates a time- or rate-dependent healing with a simi-
 511 lar order of magnitude as the healing rate b . In addition, we observe a decrease in av-
 512 erage frictional strength $\bar{\mu}$ with increasing loading rate v_L (Fig. 8b). The average slope
 513 of all four stiffnesses is negative ranging from $(\alpha - \beta) = -0.0027$ to $(\alpha - \beta) = -0.0067$
 514 indicating velocity weakening conditions. The extremely low error for the individual fits
 515 is the result of the large amount of data points for each experiment (≈ 10 million) be-
 516 cause the complete time series is used for fitting. This drastically narrows the confidence
 517 band for the slope.

518 Furthermore, comparing the reloading time t_{rel} with the loading velocity v_L we find
 519 a power-law dependency with exponents $B > -1$ (Fig. 8c). This shows a longer reload-
 520 ing time than extrapolated for the simple increase in loading velocity which is another
 521 indicator for time- or rate-dependent healing.

522 4 Discussion

523 4.1 Similarity of Constitutive Parameters

524 In rate and state friction three key parameters are determined, the direct effect a ,
 525 the healing effect b , and the state evolution variable ϕ (J. H. Dieterich, 1979a; Marone,
 526 1998). From our type of experiments we can not observe the evolution of friction directly
 527 because our system is inherently unstable. This is due to the system stiffness k_M which
 528 is below the critical stiffness k_C .

529 The healing rate $b = 0.0057 \pm 0.0005$ which is equivalent to a frictional strength-
 530 ening rate $\beta = 0.0122 \pm 0.0005$ in \log_{10} -space is at the upper estimate of natural faults
 531 and fault rocks (e.g. Alpine Fault or Scheggia Fault in Carpenter et al. (2016) or other
 532 data in Marone et al. (1990); Marone (1998)). This means that the analogue fault ma-
 533 terial shows a similar amount of time-dependent healing that is observed for natural sam-
 534 ples in rock mechanical tests. The underlying physical process is different for analogue

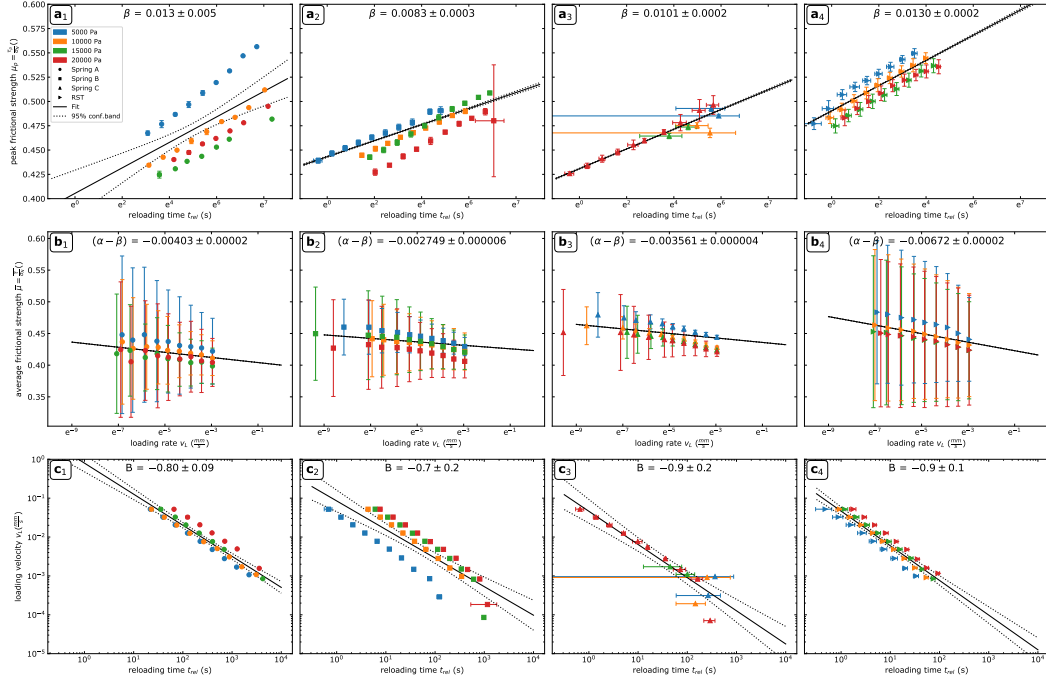


Figure 8. a) Change in peak stress with longer reloading time, which is the time between large slip events. The slope β of the log-linear fit is similar to the healing rate b from SHS-tests. b) Change in average frictional strength depending on loading rate. The slope α is an approximation of the rate-and-state parameter ($a - b$). c) Loading velocity v_L compared to reloading time t_{rel} . A negative power-law coefficient that is larger than -1 highlights longer reloading times than normal. The legend in a_1) applies to all subplots, the confidence band is derived from the covariance of fit ($2\sigma = 2\sqrt{s^2}$).

535 materials, although a certain amount of granular mechanical strengthening due to grain
 536 rearrangement is probably also present in a dry natural fault. Therefore, the glass beads
 537 are found to be usable for small-scale seismotectonic models under analogue modelling
 538 conditions. The change of frictional strength over time and average frictional strength
 539 are similar to a natural fault and can be used to simulate seismic cycles with dynamic
 540 similarity. Due to the higher healing, which is $\approx 30\%$ higher than most rocks, the ana-
 541 logue seismic cycles can be shorter in comparison with natural examples in order to rep-
 542 resent a scaled model.

543 We observe a negative direct effect $a = -0.0074 \pm 0.0031$ which is not realistic
 544 but is needed to match the value of $(b-a) = 0.0087 \pm 0.0029$ at such high healing rates.
 545 Furthermore, the friction during hold is higher ($\frac{\mu_h}{\mu_0} > 1$) just after a hold phase starts
 546 that indicates $a < 0$ because for $a > 0$ and $\frac{v_L}{v_0} < 1$ we would expect that $\frac{\mu}{\mu_0} < 1$. Be-
 547 cause of the presence of creep in the granular shear zone we assume that initially the load-
 548 ing rate is not zero but very small so that $\frac{v_L}{v_0} \ll 1$ (see also Sec. 4.2) so we could see
 549 an effect similar to a large negative velocity step. This is also due to finite machine stiff-
 550 ness k_M which was also observed by Marone and Saffer (2015). The value of $(b-a) <$
 551 0 indicates velocity weakening which results in instability under our conditions that is
 552 confirmed by the stick-slip cycles in the other experiments. Direct fitting of velocity step-
 553 ping data to get clearer results was not possible because the machine stiffness was not
 554 high enough to produce real steady state slip at our conditions.

555 Assuming that the change in peak frictional strength μ_p with increasing reloading
 556 time t_{rel} (Fig. 8) is similar to the healing rate b we find that in the main experiments
 557 the healing rate $b \approx 0.0111 \pm 0.0011$ and $(a-b) \approx -0.004266 \pm 0.000007$ which yields a
 558 direct effect $a \approx 0.0068 \pm 0.0011$. These values are in the same order of magnitude as
 559 the other estimates but show a positive direct effect. These observations however include
 560 creep and transient slip events during the reloading phase which influences the estimate
 561 of b . The estimated weakening $(a-b)$ additionally includes the effect of the stick-slip
 562 cycles which distort the calculation of the mean friction. Overall, the values are in the
 563 same order of magnitude as the estimates from our SHS-tests and agree well with the
 564 above literature values.

565 In addition, by qualitatively matching rate-and-state parameters to our SHS-tests
 566 we find that the critical slip distance D_c is in the order of 10^{-1} mm. Direct fits of the
 567 SHS-Tests obtained a $D_c \approx 200 \mu m$, by assuming an extremely low loading velocity ($v_L =$
 568 $10^{-324} \frac{mm}{s}$, smallest float represented in NumPy) during hold to obtain valid results dur-
 569 ing the hold phase. But because the results for the other parameters a and b by direct
 570 fitting using non-linear least squares were not stable, these approximations are not sta-
 571 tistically sound. Nevertheless, these values are reasonably close to values found in rock
 572 mechanical tests which vary from 2 to 100 μm (J. Dieterich, 2007, and references therein).

573 There is no statistically significant difference in the estimate of $(b-a)$ from soft
 574 and stiff systems, as expected for a material property. The scaling of strength at the on-
 575 set of slip is consistent with the findings of Beeler et al. (2001) who show the same type
 576 of scaling. The scaling coefficient typically attributed to natural rocks or gouge in the
 577 seismogenic zone, is in the same range (0.011 to 0.015 (Beeler et al., 2001); ≈ 0.01 (Scholz,
 578 1998); 0.001 to 0.01 (J. Dieterich, 2007)). Other analog model studies have used $(b -$
 579 $a)$ values in the same range to model seismotectonic processes with other materials (gel
 580 on sand paper: 0.028 (Corbi et al., 2013); rice: 0.015 (Rosenau & Oncken, 2009); cacao,
 581 ground coffee, and others: (Rosenau et al., 2017)). Therefore, we consider our models
 582 to be dynamically similar to the natural prototype, to rock deformation experiments in
 583 the MPa-range (e.g. Tullis & Weeks, 1986), and to numerical simulations of rate and state
 584 friction (e.g. Ferdowsi et al., 2013).

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4.2 State Evolution During Hold Phases

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During a hold phase the state θ of the system changes according to a certain relationship (section A2) which leads to a change in frictional strength of the fault. Beeler et al. (1994) state that purely time-dependent healing, which is given by the Aging law (Eq. A2), is independent of stiffness while the Slip law (Eq. A3) shows a dependency on stiffness because it requires active fault slip during healing. For our SHS-experiments we did not systematically vary stiffness but a change with normal stress was observed. The experiments at lowest normal stress show a higher than average healing rate. Additionally, these experiments show a decrease in stress during the hold phase. We attribute this to enhanced creep which is promoted by the low normal stress. As a result, the healing is amplified by larger amounts of slip in our experiments. In terms of constitutive laws this would mean that our material is better characterized by the Slip law than by the Aging law. A simple experimental test is to use additional data from unstressed SHS-tests (Marone, 1998). However, preliminary experiments with unloading showed that it is technically not feasible to do unstressed tests with our testing apparatus because the loading and unloading of the samples take too much time in comparison to the hold durations. We observed less healing for these tests but the dataset is very noisy so the results are statistically not relevant.

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For most experiments we observe an upwards step in stress immediately after the onset of holding which indicates a negative a which is also evident from the estimates of $(b-a)$. The step is followed by an exponential decay to a lower residual stress which shows a decay rate $\lambda = 2.1 \pm 1.2 s^{-1}$. The stress decays to the residual stress after hold within less than 3 seconds ($< 1\%$ difference). This indicates that creep due to shear stress quickly dissipates and the sample is not slipping along the shear surface during hold. The thickness of the granular packages first oscillates at a frequency of around 19Hz and stabilizes to a constant value within the same time as the shear stress. These observations indicate that there is no measurable slip during hold which means that slip during hold is minute and therefore the Aging law is more appropriate. In spite of many points that speak for the Aging law, according to Bhattacharya et al. (2017) it is not sufficient to only look at the evolution of stress during hold phases but also to model the experimental data numerically. Therefore, we cannot exclude the possibility that other state evolution laws apply and additional experiments with different stiffnesses and numerical modelling of the actual data are needed to clarify this finding.

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4.3 Micromechanical processes

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Granular material gains shear strength due to force chains oriented in the direction of the maximum stress (Cates et al., 1998). Depending on the number, length and orientation distribution of such chains shear deformation might be stable or unstable. Stick-slip is therefore interpreted as a cyclic setup and breakdown of force chains, the frequency and size of which should be a function of grain size distribution (Mair et al., 2002). Furthermore, granular materials exhibit so called 'jammed states', where jamming is induced at high packaging density or by application of shear stress (Bi et al., 2011). We corroborate this view as large slip events are associated with compaction while the interseismic period is characterized by accelerating creep and dilation (Figure 2).

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The normal stress is one of the critical factors that control the creep threshold of the system. For low normal stresses it is easier for the grains to rearrange during the creep phase. Firstly, this results in higher background slip of grains that exhibit a much lower normal stress along their contacts and can easily slide along each other. Secondly, the ratio of normal stress to dilatational stress, that pushes the grains apart when sliding over the rough internal shear zone, is smaller. Therefore, the force chains are less effective in strengthening the material at low confining pressures.

635 The occurrence of small slip events is in accordance with other studies that show
 636 transient effects during the transition of the stick phase to dynamic slip (Nasuno et al.,
 637 1998; Ferdowsi et al., 2013). Because they are much smaller than the main events it is
 638 suggested that the events are the expression of internal reorganization in the granular
 639 material. During this internal deformation the grains are jammed and the force chains
 640 are rearranged into a more stable configuration. Although creep continues the newly formed
 641 granular package is stronger than the previous package and therefore a short period of
 642 quiescence without slip events occurs. This rearrangement can occur several times dur-
 643 ing the late interevent phase. If the internal structure reaches a critical threshold, prob-
 644 ably determined by the contact ratio and packing density, a runoff process starts and the
 645 system changes from creeping to dynamical slip.

646 Other studies have shown a similar system behaviour that is attributed to inter-
 647 mittent criticality (Ben-Zion et al., 2003). In contrast, to the self-organized critical sys-
 648 tem, intermittent criticality implies a cyclic evolution of the fault zone, whereas the SOC
 649 only gives a general statistic fluctuation around the critical state (i.e. failure criterion).
 650 If we apply the concept of intermittent criticality, the small precursors are the expres-
 651 sion of small scale stress perturbations along the fault zone. Overall the stress field within
 652 the granular fault zone homogenizes by increasing rearrangement of force chains, that
 653 explains the increasing frequency of events up to a certain point. Then the system is largely
 654 homogenized and is in a critical state, very close to failure, which is comparable with the
 655 state of stress in the lithosphere (Sornette et al., 1990). This behaviour has also been
 656 observed for the temporal and spatial clustering of smaller earthquakes (Hainzl, 2003).

657 The behaviour of dilation during the interevent cycle is even more complex and it
 658 is difficult to assign a direct relation to micromechanical processes. The observed increase
 659 in wavelength of the small amplitude oscillations could indicate a smoothing of the in-
 660 ternal fault surface, leading to a smoother frictional response. The discrete upward and
 661 downward steps might be artificial, or the result of sensor noise. However, the strong re-
 662 producibility over multiple cycles indicates that mechanical explanations can be valid,
 663 too. For example, internal reorganization of the granular packaging leads to discrete con-
 664 formations of packaging with different densities that are characteristic for each state of
 665 the system.

666 4.4 Slip Modes

667 4.4.1 Criticality of Analogue Fault

668 From the determined rate-and-state parameters we can now derive the critical stiff-
 669 ness to evaluate how close the main series experiments are to the bifurcation from sta-
 670 ble and unstable slip. Because normal stress is constant in each series we use a stabil-
 671 ity criteria according to J. Dieterich (2007):

$$672 \quad k_c = \frac{\zeta \sigma_N}{D_c} \quad (7)$$

673 According to Heslot et al. (1994) the critical stiffness normalized by normal load
 674 σ_N (in their case slider mass M) is a slowly decreasing function depending on loading
 675 velocity v_L . Accordingly, we correct for loading rate v_L with respect to the loading rate
 of the SHS-tests $v_0 = 0.52 \frac{mm}{s}$ and the scaling factor $\alpha = 10$:

$$676 \quad k(V) = k_c - \alpha \ln \left(\frac{v_L}{v_0} \right) \quad (8)$$

677 Utilizing both estimates for $\zeta = (b - a) = \{0.0084, 0.0130\}$, three different es-
 timates for $D_c = \{50, 100, 200\} \mu m$ and normalizing by normal stress the normalized

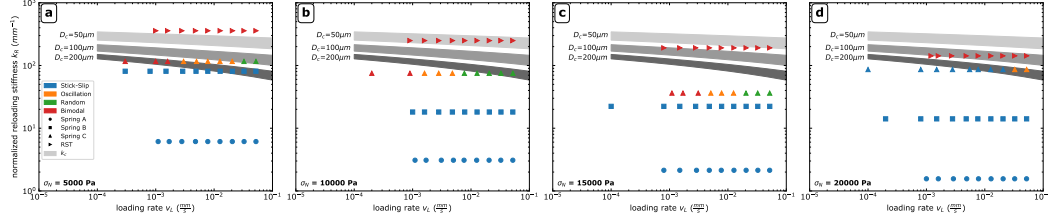


Figure 9. Slip modes in the $k - v$ space represented by the normalized reloading stiffness which includes the material’s effect. The critical stiffness k_c is calculated from the rate-and-state parameters from SHS-experiments at maximum stiffness (Section 3.5) and explains the transition from unstable to stable (random) slip mode for Spring C.

678 critical stiffness k_c ranges between 60 and 340 mm^{-1} . Comparing this with the normal-
 679 ized machine stiffness k_N (Fig. 3) we find that most experiments show $k > k_c$ and there-
 680 fore should only show stable sliding (J. Dieterich, 2007). However, due to the material
 681 inside the machine the actual stiffness of the system is lower. If the normalized reload-
 682 ing stiffness k_R is used instead (Fig. 9), the experiments with lowest stiffness now show
 683 $k \leq k_c$ and the experiments with higher stiffness now show $k \approx k_c$.

684 Consequently, the experiments with Spring C (Δ in Fig. 9) now fit with the change
 685 of unstable to stable slip when transitioning the stability boundary (grey boxes in Fig.
 686 9). It is unclear why the stiffest experiments (RST) still shows unstable slip. Using a higher
 687 scaling factor $\alpha = 50$ shifts the stability boundary upwards. Therefore we suspect that
 688 the scaling factor α might represent the ratio of normalized machine stiffness k_N to nor-
 689 malized reloading stiffness k_R which is $\frac{k_R}{k_N} \approx 50$ for RST and $\frac{k_R}{k_N} \approx 10$ for the experi-
 690 ments with a spring (Fig. 4). From these observations we also think that it is safe to as-
 691 sume that D_c is in the order of 100 to 200 μm which is roughly the radius of a glass bead
 692 $r_{GB} = 150 \dots 200 \mu m$. Due to the uncertainties in the estimation of k_N and the high
 693 uncertainty for D_c the values for k_c and the location of the stability boundary are not
 694 well constrained. The uncertainty for D_c might result from the possibility that D_c is not
 695 constant because the thickness of the active shear zone might change during the exper-
 696 iment which is a primary factor for the scaling of D_c (Marone & Kilgore, 1993). Further-
 697 more D_c shows a dependency on slip velocity which could not be studied with our setup
 698 (Hatano, 2009). Nevertheless, the fits of the SHS tests and main experiments yield re-
 699 sults that seem valid and the stability boundary lies within the same order of magnitude.

700 4.4.2 Modelling of Slow Slip Events

701 For natural examples the slip mode is usually identified by the relation of seismic
 702 moment M_0 and characteristic duration T (e.g. Ide et al., 2007; Gombert et al., 2016).
 703 Regular earthquakes show a much shorter duration ($M_0 \propto T^3$) in comparison with slow
 704 earthquakes ($M_0 \propto T$). The scaling relation leads to a characteristic separation between
 705 events of equivalent seismic moment M_0 . We observe a similar separation of slow and
 706 fast events which is most prominent for the experiments at highest stiffness. Depend-
 707 ing on the actual seismic moment the separation in nature is between 2 and 5 orders of
 708 magnitude. In the experiments at highest stiffness we observe a separation in average
 709 slip rate of 2-3 orders of magnitude. The difference in peak slip rate is up to 5 orders
 710 of magnitude. This separation is smaller but still statistically significant for lower ma-
 711 chine stiffness where we also find oscillating slip modes. Similarly, the frictional stress
 712 drop that can be seen as a proxy for seismic moment in our experiments, shows a sep-
 713 aration of 2 orders of magnitude which indicates that the dynamics are different for slow
 714 and fast events.

715 Several studies highlight the relationship of transient slip events promoting seis-
716 mic activity (A. Kato & Ben-Zion, 2021, and references therein). The results from our
717 study suggest that glass beads as analogue fault gouge shows a large variety of slip modes
718 not only at high mean stresses as previously found by others (J. H. Dieterich & Kilgore,
719 1996; Cain et al., 2001; Mair et al., 2002; Cui et al., 2016). Therefore, the usage of glass
720 beads in small scale analogue models with intermediate to high stiffness at stresses in
721 the kPa range is suitable to model seismic fault behaviour. For example the glass beads
722 can be used as fault gouge in between two elastic blocks in gel-slider type models (similar
723 to Corbi et al., 2011) or in more complex models. If the stiffness of the model is adjusted
724 correctly (e.g. in the range of Spring C) several slip modes are possible depending on
725 the normal stress on the fault. The normal stress regime on the fault can then be de-
726 signed by geometric orientation of the fault with respect to the loading direction. More
727 complex fault geometries, possibly forming fault networks, then lead to transient stress
728 on the individual faults and thereby altering slip mode. These setups can be used to study
729 the complex interplay of fault geometry and slip on individual faults, while retaining dy-
730 namic and kinematic similarity. Due to transient stress changes individual analogue faults
731 might change their slip mode from pure stick-slip to creep and thereby changing system
732 dynamics and the activity of other faults in the system.

733 Especially the temporal distribution of slow events between the occurrence of fast
734 events highlights the possible application of glass beads as analogue fault gouge. The in-
735 creased probability of slow events towards the end of the reloading time and the high
736 stress level is similar to the behaviour observed for large fault systems that show increased
737 seismic activity towards the end of the seismic cycle (A. Kato & Ben-Zion, 2021). The
738 abrupt decrease in probability just before failure might be an expression of strong lock-
739 ing due to jamming by shear (Bi et al., 2011) and might favour fast slip events by driv-
740 ing the frictional strength above a critical threshold. In general we find that the 'pre-
741 cursory' phase where the fault shows signs of imminent failure starts relatively early, with
742 the onset of creep at about 50% of the reloading time. This is similar to long prepara-
743 tory phases of large earthquakes (Bouchon et al., 2013) and to recent findings by Igarashi
744 and Kato (2021). In our case the slow events do not always act as precursors because
745 for certain conditions they occur in a repeating pattern at high stresses (e.g. Spring C
746 at low normal stress) and therefore are reminiscent of 'similar earthquakes' (Igarashi &
747 Kato, 2021). Although the stress drop magnitude slowly increases over several repeat-
748 ing events, they do not show a clear threshold for which the slow event grows into a large
749 event. Consequently, they only pinpoint that the fault is close to failure but not that the
750 fault will fail with a fast slip event after a certain type of slow event.

751 The slip behaviour of the granular analogue is the result of the interaction of fric-
752 tion with the complex network of force chains that is created in the sheared bulk ma-
753 terial (Cates et al., 1998; Daniels & Hayman, 2008). This micromechanical mechanism
754 creates the macroscopic behaviour that can be described with rate-and-state friction. Our
755 findings support the hypothesis that rate-and-state like dynamics are the expression of
756 processes that emerge close to the criticality boundary. Similar kinematic observations
757 can be made for a range of microphysical processes and conditions (e.g. Kabla et al., 2005;
758 Papanikolaou et al., 2013; Scuderi et al., 2015; Hecke, 2009; Lemaître & Caroli, 2009).
759 Denisov et al. (2016) show that force fluctuations in granular matter, of which our ex-
760 periments but also earthquakes are a part of, show scaling relations that are akin to crit-
761 ical phenomena. Furthermore, the size and stress distribution of slip events within gran-
762 ular materials at the same conditions as in our experiments is found to be universally
763 related with slip events on multiple scales and follows similar scaling relations (Uhl et
764 al., 2015). The underlying physical mechanisms for these observations can be quite dif-
765 ferent but yielding the same results on a macroscopic scale which is similar to the de-
766 scription of flows using rheological equations that purely describe the observed relation
767 of motion and stresses while the actual deformation mechanism is different in different

768 fluids. Therefore, we find that our model has widespread possibility of application within
 769 seismotectonics, engineering and hazard assessment for earthquakes and landslides.

770 5 Conclusions

771 We have used an annular shear apparatus to characterize the stick-slip behaviour
 772 of a granular fault zone analogue composed of glass beads. Using slide-hold-slide tests
 773 the rate-and-state properties have been qualitatively evaluated and quantified. The heal-
 774 ing rate is found to be $b = 0.0057 \pm 0.0005$. The direct effect a is quantified by two
 775 approaches and found to be $a = -0.0076$ for estimates from the change of peak stress
 776 with increasing reloading velocity in SHS-tests, while the procedure by Beeler et al. (2001)
 777 with a specific $\frac{v_L}{t_h}$ -ratio yields $a = -0.0030 \pm 0.0030$. In both cases the material is found
 778 to be velocity weakening with a $(b - a) = 0.0131 \pm 0.0031$ or $(b - a) = 0.0087 \pm 0.0029$
 779 respectively. Due to the evolution of stress during the hold phase we find the Aging law
 780 to be slightly more appropriate for our material but the results from SHS-tests only re-
 781 main inconclusive. The critical slip distance D_c is estimated to be in the sub-mm range
 782 but can not be quantified with the presented setup. The effect of machine stiffness k_M ,
 783 loading rate v_L and normal stress σ_N on the slip mode is studied. We find a large va-
 784 riety of slip modes ranging from pure stick-slip, oscillations to bimodal slip modes within
 785 the same experiment by only varying certain extrinsic parameters which fits well with
 786 the stability boundary derived from the rate and state parameters. Low stiffness, low
 787 loading rates and high normal stresses favour pure stick-slip with small amounts of in-
 788 terevent creep. Higher stiffness, especially in combination with low loading rates, leads
 789 to a bimodal distribution of fast, large events that are preceded by slow, small events.
 790 The slip events reproduce typical characteristics that have been observed in similar ex-
 791 periments in other experimental setups with different boundary conditions and mate-
 792 rials allowing to generalize the observations to natural occurrences of earthquakes. In
 793 the experiments, rearrangement in the granular package is the major micromechanical
 794 process which distributes and dissipates stress during shear. This drives the system closer
 795 to criticality leading to the observed precursory strengthening and the short period of
 796 quiescence before a large slip event. We conclude that the small transients can strongly
 797 affect the statistical characteristics of a single fault zone system and makes the mate-
 798 rial suitable for the use in larger analogue modelling setups that model seismotectonic
 799 deformation with a higher geometrical complexity. The small scale events during the pre-
 800 cursory phase are the expression of distributed fluctuations of the system in a critical
 801 state. Further examination of these fluctuations and their correlation with the genera-
 802 tion of large events may give important constraints on the predictability of slip events
 803 (as suggested by Ben-Zion et al., 2003). Furthermore, the higher complexity with dif-
 804 fering slip modes due to the characteristics of the glass beads could provide additional
 805 insights into the system behaviour and the interaction of faults in analogue models that
 806 are closer to the behaviour of a natural fault zone. The results from this study shed light
 807 on the micromechanical mechanisms from which rate and state-friction emerges. There-
 808 fore it can act as a benchmark for numerical models of fault zones, alleviate the design
 809 of more complex analogue models and helps interpreting kinematic natural observations
 810 of fault slip.

811 Appendix A Data Analysis

812 The experimental data is examined using a combination of classical event detec-
 813 tion, statistics and machine learning. To analyze the occurrence and properties of the
 814 slip events we employ a peak detection that is based on a minimum stress drop thresh-
 815 old for each experiment. Then we extract certain characteristic points in the stress curve,
 816 these are highlighted in Fig. 2c. To avoid confusion with other terms which might de-
 817 scribe similar features Tab. A1 lists all the characteristics we use and highlights pub-
 818 lications where a more throughout definition is found.

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A1 Picking and First Order Properties

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For first order characterization of the experiments we use a simple peak detection to find slip events in the stress curve. For this the data is split into sets of equal loading rate, normal stress and stiffness. A fixed threshold for stress drop per set facilitates the detection of sudden changes in shear stress. Fine tuning this value enables the detection of large and fast, but also of small and slow events by searching for positive and negative peaks in the data. The result of peak detection is cross checked by manual inspection of the stress curves and the detected peaks (Fig. 2a+b).

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The point X of maximum stress immediately before failure is denoted by X_p indicating peak values, the point of minimal stress after a slip event is indicated by X_e accordingly (Fig. 2c). In most cases X is replaced by the appropriate physical quantity such as shear stress τ or velocity v . A velocity threshold defines the separation of non-dynamic and dynamic slip events. A slip event is considered dynamic when at any point during a decrease of shear stress the slip velocity v_s is higher than the threshold v_d . In this study we used the maximum loading velocity of $v_L = 0.02 \frac{mm}{s}$ as the threshold. This allows the definition of onset of dynamic slip, denoted by X_d , maximum slip X_m where slip velocity is at its maximum and the end of dynamic slip X_f where the slip velocity drops below the critical value. These points now define a full cycle, which we see as an analogue of a seismic cycle.

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The full cycle is defined as the period of time between two slip events that have a dynamic phase with velocities above the threshold v_d . During the majority of a full cycle the shear stress increases in a linear relation with load point displacement but deviates to a non-linear relation. This point is the onset of creep and is defined as the point where the linear trend extrapolated from the previous points deviates by more than 1%. The slope of the linear trend, calculated by least squares fitting, also defines the cyclic reloading stiffness k_L which is a measure for the overall stiffness of the setup including the bulk stiffness of the granular material. This stiffness is also used for further calculations of the criticality using the rate-and-state framework.

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Assuming an overall elastic behaviour of the granular material when completely locked, we can estimate the amount of creep either as overall proportion or as instantaneous creep. Overall creep is calculated by linearly extrapolating the shear stress increase over the full cycle using k_L as a slope and then relating the predicted and observed point of failure. Similarly the instantaneous creep is calculated by a similar method but doing a point wise calculation.

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A2 The RSD-Formulation

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Following J. Dieterich (2007, and references therein), shear stress τ evolves as a function of effective normal stress σ , load point velocity v_L and a set of experimentally derived parameters μ_0 , a , b , θ , D_c in relation to a reference load point velocity v_L^* :

$$\tau = \sigma \left[\mu_0 + a \ln \left(\frac{v_L}{v_L^*} \right) + b \left(\frac{\theta v_L^*}{D_c} \right) \right] \quad (\text{A1})$$

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This is a heuristic description of the change in shear stress in response to a change in slip velocity. The parameters in Eq. A1 are usually derived experimentally using velocity stepping tests where the system sliding at a given reference load point velocity v_L^* under stable conditions ($\dot{\theta} = 0$) is perturbed by setting a new loading velocity v_L . This prompts an direct reaction of shear stress that is proportional to the magnitude of the perturbation $\frac{v_L}{v_L^*}$ and the constant a ('direct effect'). Following this immediate reaction, shear stress adjusts to a new level defined by the evolution of state over time in relation to the new loading velocity normalized by the characteristic slip distance $\frac{\theta v_L^*}{D_c}$ and a con-

stant b ('evolution effect'). The evolution of state over time $\dot{\theta}$ is defined by choosing one of the following evolution laws:

$$\dot{\theta} = 1 - \frac{v_L \theta}{D_c} \text{ Aging Law} \quad (\text{A2})$$

$$\dot{\theta} = -\frac{v_L \theta}{D_c} \ln \frac{v_L \theta}{D_c} \text{ Slip Law} \quad (\text{A3})$$

$$\dot{\theta} = e^{\frac{v_L}{\delta c}} - \frac{v_L \theta}{D_c} \ln \frac{v_L \theta}{D_c} \text{ Kato Law} \quad (\text{A4})$$

$$\dot{\theta} = 1 - \frac{v_L \theta}{D_c} - \frac{c}{b} \theta \frac{\dot{\tau}}{\sigma} \text{ Nagata Law} \quad (\text{A5})$$

The Aging law (Dieterich, 1978) and Slip law (Ruina, 1983) are the most commonly used state equations, while the Kato law (N. Kato & Tullis, 2001) and Nagata law (Nagata et al., 2012) are more recent developments. All of the above laws contain a slip dependent component which is expressed in the term $\frac{v_L \theta}{D_c}$. Consequently, the Slip law (Eq. A3) does not show any healing when there is no slip ($v_L \rightarrow 0$) which can be tested through unstressed SHS tests where the sample at rest is not under stress and thus slip along grain contacts is hindered. The Aging law (Eq. A2) shows constant healing at rest due to the 1 in the equation. This purely time-dependent effect makes the Aging law fit better to experimental data (Beeler et al., 1994). However, for large velocity steps ($|\log_{10} \frac{v_L}{v_L^*}| > 2$) the Aging law shows a linear decay that is dependent on the sign and magnitude of the step which is not in accordance with experimental data. Furthermore, the Aging law does not fit well to the state evolution during a hold phase and needs non-constant a , b and D_c (Bhattacharya et al., 2017). The Slip law exhibits a better fit for large velocity steps and for the evolution of state during a hold. This resulted in a reformulated version of the Slip law that accounted for time-dependent healing by N. Kato and Tullis (2001) termed Kato law by Bhattacharya et al. (2017) (Eq. A4). A further improvement to the previous laws has been proposed by Nagata et al. (2012) which incorporates the relation of the 'evolution effect' b and a new constant c , as well as the normalized stressing rate $\frac{\dot{\tau}}{\sigma}$ (Eq. A5).

A21 Tests and Derived Quantities

We performed SHS tests with stressed hold phases (Marone, 1998) to determine the direct effect a , rate of healing b and the appropriate state law.

Due to the evolution of state during a hold, the frictional resistance μ_p of a granular medium increases with the natural logarithm of the hold time t_h and gives rise to the healing rate b (Bhattacharya et al., 2017):

$$b = \frac{\delta \Delta \mu_p}{\delta \ln t_h} \quad (\text{A6})$$

This increase in μ_p is measured with slide-hold-slide tests. During the first slide phase a steady-state value of μ_s is established. Then the machine is stopped for a certain time t_h , either under stress or unstressed. Then the sample is resheared and the increase of peak strength with respect to the previously established stable sliding resistance is measured as $\Delta \mu_p = \mu_p - \mu_s$. Relating the results to $\ln t_h$ a linear increase can be measured, which is the healing rate b . Furthermore, the loading velocity v_L plays an important role in the magnitude of frictional resistance after loading μ_p but should not in-

fluence the healing rate b (Beeler et al., 2001). Experiments at increasing v_L therefore show the same slope b but increasing $\Delta\tau_p$.

This effect can be exploited to determine the direct effect a from slide-hold-slide tests by using a specific spacing between the realized loading rates (Fig. 9 in Beeler et al., 2001). If the ratio of the loading velocities $\frac{v_{L1}}{v_{L2}}$ is equal to the ratio of hold times $\frac{t_{h2}}{t_{h1}}$ (Eq. A7) then the increase in $\Delta\mu_p$ is proportional to $a \cdot \ln \frac{v_{L1}}{v_{L2}}$. We estimated the direct effect a from the average increase in $\Delta\mu_p$ over all realized hold times using a set of SHS-Tests that fulfil eq. A7.

$$\frac{v_{L1}}{v_{L2}} = \frac{t_{h2}}{t_{h1}} \quad (\text{A7})$$

Furthermore, the direct effect can be measured from the offset of the y-intersect of individual fits of the SHS tests at different velocities by:

$$a = \frac{y_0 - y_1}{\ln \frac{v_0}{v_1}} \quad (\text{A8})$$

Other approaches to estimate $(a-b)$ were also tested with our data. Corbi et al. (2013) defines $(a-b)$ using the peak friction μ_p and sliding velocity v_L (Eq. A9). For this we determine the frictional resistance μ_p at the peak point τ_p just before a dynamic failure because at the peak, there is a plateau of shear stress and therefore the current slip rate of the fault equals the loading rate $v_S = v_L$.

$$(a - b) = \frac{\Delta\mu_p}{\Delta \ln v_L} \quad (\text{A9})$$

914 **A3 Supporting Figures and Tables**

Table A1. Terminology and definition of characteristic points.

Term	Symbol	Definition
Stiffness	k_M	Theoretical stiffness of machine calculated from spring stiffness and apparatus stiffness.
Cyclic reloading stiffness	k_L	Measured relationship of force increase and load point displacement during an interevent phase. This is approximately the real stiffness of the apparatus with granular material.
Unloading stiffness	k_U	Measured relationship of force drop and horizontal lid displacement during a slip event. This quantity is measured using a high speed camera for each of the realized stiffnesses.
Loading velocity	v_L	Rotation velocity of the shear cell during an experiment. This value is defined as the velocity along the median circumference of the shear cell which divides the cell area into two equal compartments.
Slip velocity	v_S	Velocity of the lid during an event along the same circumference as the loading velocity. Calculated from F_D and k_U .
Load point displacement	d_L	Horizontal displacement of the shear cell along the median circumference by the loading velocity. Calculated by integrating v_L over time.
Slip displacement	d_S	Horizontal displacement of the lid during a slip event. Calculated by integrating v_S .
Lid displacement	d_H	Vertical displacement of the lid due to internal deformation of the granular medium. The zero-level is defined as the top of the shear cell.
Package density	ρ_P	Density of granular material during the experiment. Calculated from weighted mass, shear cell area and d_H .
Slip event		Abrupt reduction in shear stress along the shear zone coinciding with a counter rotation of the lid. Has a start ‘Event start’ and end ‘Event end’ defined by characteristic points on the shear curve.
Microslip		Similar to a ‘Slip event’ but with intensity and slip velocity a few orders of magnitude lower.
Recurrence time	t_r	Time between the end (t_e) of a ‘Slip event’ and the start of the next (t_p). The time span is named interevent phase analogous to the interseismic period for earthquakes.
Event peak	p	Maximum shear stress before a ‘Slip event’.
Event end	e	Minimum shear stress after a ‘Slip event’.
Onset of dynamic event	d	Critical point where slip velocity during an event is larger than the loading rate.
End of dynamic event	f	Critical point where slip velocity during an event is lower than the loading rate.
Preslip	d_p	Slip that happens during the acceleration of a slip event between ‘Event peak’ and ‘Onset of dynamic event’.
Creep		Ratio of ‘Slip displacement’ and ‘Load point displacement’ during the interseismic phase. Due to permanent internal deformation at very low rates, there is a deficit between displacement that is imposed on the sample and released slip during a ‘Slip event’.
Onset of Creep	c	Position on the shear stress curve where the reloading deviates from the linear trend (defined as ‘cyclic reloading stiffness k_L ’) by more than 1%.

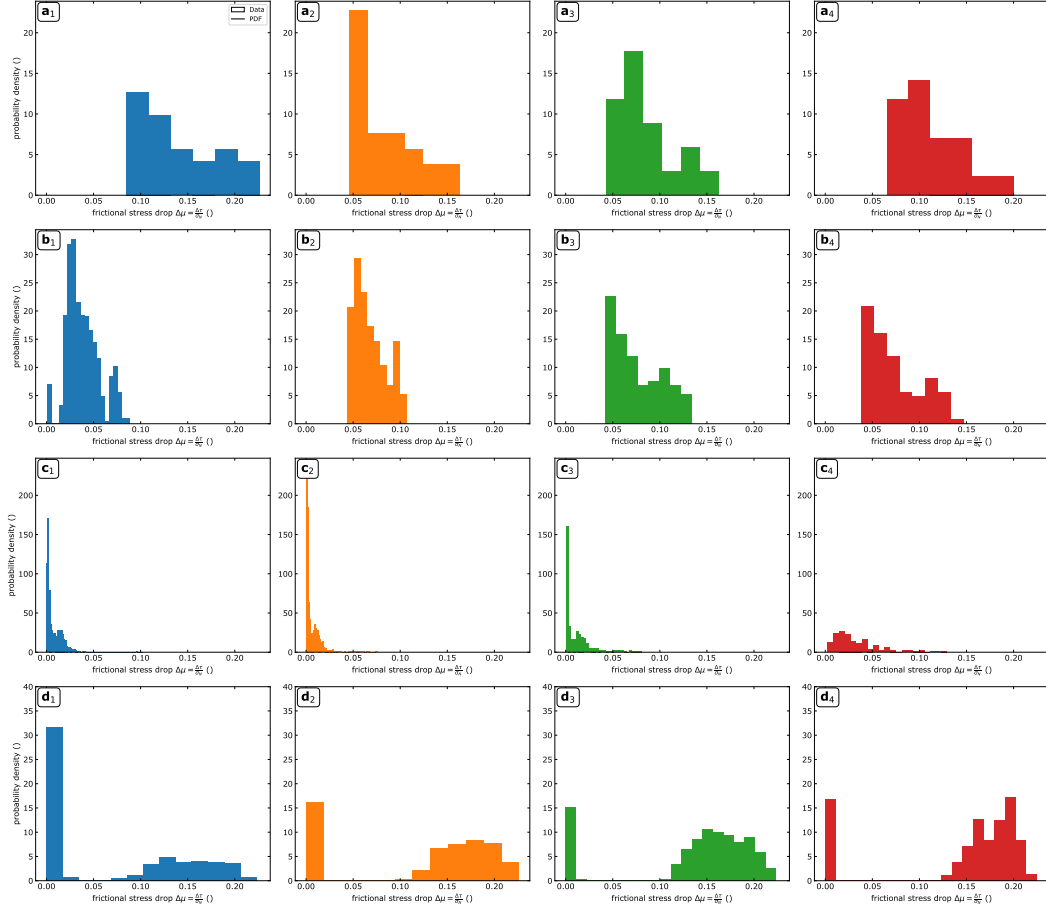


Figure A1. Histograms for frictional stress drop per experiment series. The legend in a_1 applies to all plots. Each subplot summarizes data from different loading rates but at constant normal stress (indicated by colour) and constant stiffness. Each row has constant stiffness with a) $k_M = 3.3 \frac{kN}{mm}$, b) $k_M = 19.6 \frac{kN}{mm}$, c) $k_M = 82.6 \frac{kN}{mm}$, and d) $k_M = 1354.0 \frac{kN}{mm}$.

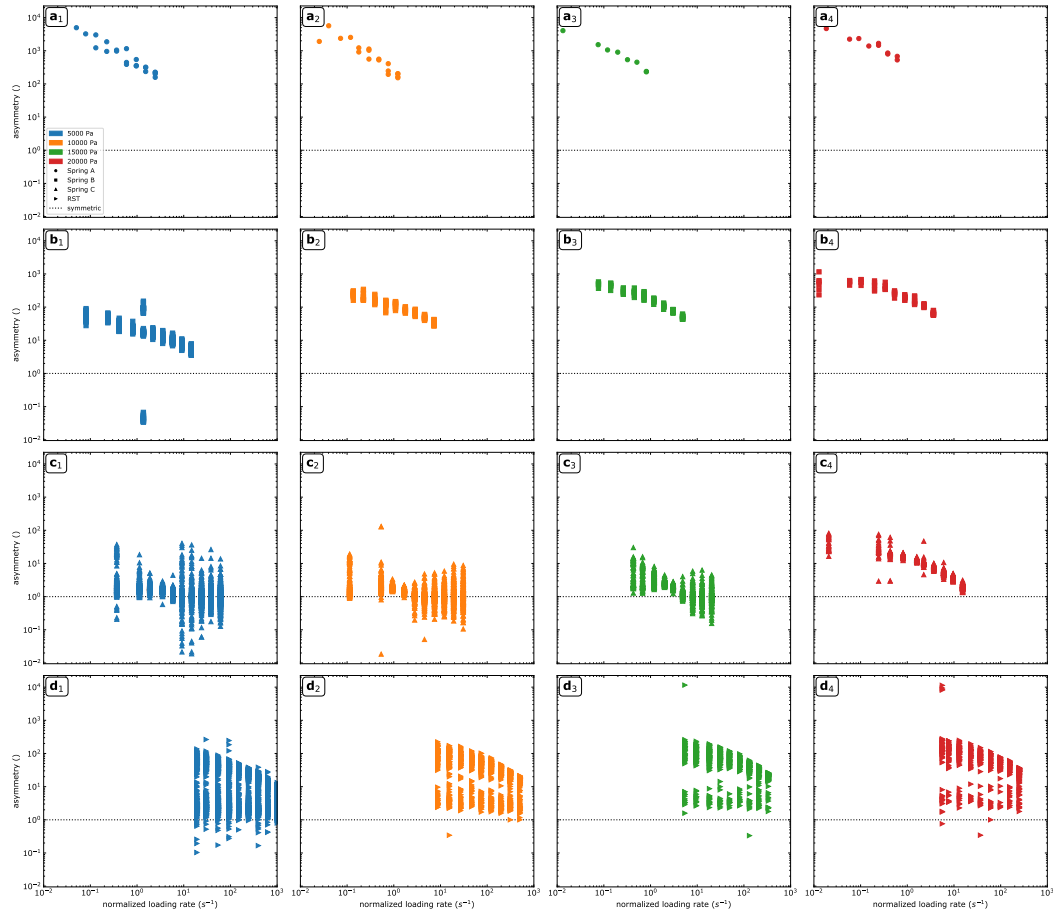


Figure A2. Asymmetry for all events and all experiments. The legend in a_1 applies for all plots. Each row represents experiments of the same stiffness which is also indicated by the individual markers. Colour highlights the different normal stresses which has an additional influence besides the stiffness and normalized loading rate.

Acknowledgments

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