NOTES ON STATISTICAL AGE DISPERSION IN FISSION-TRACK DATASETS: THE CHI-SQUARE TEST, ANNEALING VARIABILITY, AND ANALYTICAL CONSIDERATIONS

Kalin T. McDannell 1,2*

¹Geological Survey of Canada, Calgary, Canada ²Dartmouth College, Hanover, NH, United States *correspondence: kalin.t.mcdannell@dartmouth.edu March 5, 2020 (edited Mar. 27, 2020 and Nov. 3, 2021)

The manner in which we apply statistical analysis is founded on an idealized concept of how fission tracks are generated and measured, which may be more complicated in practice when there are departures from ideal circumstances. The fission track (FT) method is a low precision technique, and we beat this shortcoming with numbers, that is to say, by repeating many low precision individual analyses to provide a better estimate of some 'true' age or to make inferences about the overall age population. The chi-square (χ^2) test is a simple diagnostic used to assess whether single-grain age estimates are statistically homogenous and consistent with a common true age [1]. Significance of the test is assessed by calculating the *p*-value, which is the probability that a value drawn from the χ^2 (n-1) distribution exceeds the χ^2 statistic [1]. A small *p*-value indicates that the data are inconsistent with a common ρ_s/ρ_i ratio for the external detector method, EDM, or a common value of $N_s/(^{238}U/^{43}Ca)$ for the apatite LA-ICP-MS ζ -method. The convention being that a *p*-value < 0.05 offers sufficient evidence against the null hypothesis (of a common true age) and a *p*-value < 0.01 is strong evidence against the null hypothesis [1]. Typically, a χ^2 value of < 0.05 is considered to indicate that the single-grain ages are unlikely to be drawn from a single Poissonian distribution with a discrete mean value and may represent multiple age populations [1–3]. However, a large *p*-value only means there is a lack of evidence against the null hypothesis—it does not lend support for an alternative hypothesis [1].

The χ^2 test applied to FT data is not a data quality filter. A χ^2 pass or failure does not equate to 'good' or 'bad' fission track data, respectively. A sample should still undergo evaluation if it passes the χ^2 test, and if a sample fails, an attempt should be made to understand why it failed. Even though the χ^2 statistic is designed to test the underlying assumption of a single, underlying age component for a sample, a problem encountered in many published fission-track studies is that the results of the χ^2 test are noted when a sample passes, but are ignored or downplayed if the sample fails—leading to a tendency to incorrectly judge the quality of FT data based on the χ^2 test. This is likely due to the fact that without more information, we do not necessarily know how to explain a χ^2 failure. For example, it has been acknowledged for decades that chemical composition has an effect on fission-track annealing and consequently FT ages [e.g., 4-8], however many studies lack proper quantification of mineral composition to evaluate potential cause(s) of high age dispersion and χ^2 failure. The amount of thermal annealing, and therefore FT shortening, experienced by a sample may differ between grains due to variable chemical composition [2, 5–9]. In this case, a distribution of FT single-grain ages should exist, rather than a single common value [1, 3]. This point is critical—if a crystalline basement or detrital fission-track sample is characterized by variable intra-grain chemical composition (or additionally, variable provenance for detrital samples) and has experienced thermal annealing—then a χ^2 failure should be expected. This sample would be considered multikinetic and if the relationship between mineral chemical composition and discrete age populations can be characterized, then each kinetic population within a sample will be sensitive to a specific temperature range (i.e. variable thermal

annealing kinetics), broadening the overall sensitivity of the sample [10–12]. This essentially means that multiple thermochronometers are contained within a single sample—characterized by kinetic age populations with distinctive annealing behavior. Multikinetic behavior is extremely useful because it allows the user to 'do more with less,' in that we can constrain more thermal information with a single sample and yield improved time-temperature resolution for thermal history analysis, which is often the goal of nearly all FT studies.

There are also other factors that are important when assessing and dealing with fission track age dispersion that require attention when interpreting a dataset. Some of the factors that play a role in statistical age dispersion include: (i) the total number of grains analyzed, (ii) the precision of individual fission track grain ages within a population, and (iii) the number of spontaneous tracks (N_s) counted by the analyst. Uranium concentration and N_s are the main contributors to the analytical uncertainty during fission track data collection [13, 14]. There is some concern that the standard error of LA-ICP-MS ages may be underestimated, causing age over-dispersion, and that the lack of a 'matched pairs' design (as with the EDM), may produce inaccurate ages due to difficulty in accounting for U heterogeneity [13]. This is more of an issue for very young or U-poor apatite grains, but ways to alleviate this are to only count tracks in the laser ablation spot or carry out multiple U spot analyses on a grain [13]. Nevertheless, a closer examination of LA-ICP-MS U-error propagation may be warranted.

Young or U-poor samples typically have low N_s and larger associated errors, whereas old or U-rich samples have high N_s and higher precision ages. Generally, if a population of grain ages is characterized by low precision, then there is a χ^2 pass (p-value > 0.05) and low overall age dispersion would be expected, whereas if a dataset contains highly precise ages, then a χ^2 failure would be expected [1, 2]. The order-of-magnitude better age precision obtained by the LA-ICP-MS fission track technique (partly due to direct and more precise measurement of U compared to the EDM) will often produce χ^2 failures that are difficult to interpret without additional information such as mineral composition, and complicates parsing of kinetic populations that align with radial plot mixture model predictions.

An exaggerated example to illustrate the effect of age errors is shown by FIGURE 1 using a radial plot [15] to graphically display single-grain ages and their precisions for two synthetic samples of 20 grains each with randomly generated ages between 50–100 Ma. We hold the grain ages fixed for both samples and only change the individual age errors at random. The first sample is precise (i.e., more ages further from the origin on the radial plot; FIGURE 1A), with individual grain-age errors varying between ~5–15%, while the second sample is low precision (FIGURE 1B) with age errors between ~12–85%. Clearly, the errors on individual ages influence the sample χ^2 probability (0.0 vs. 0.98), age dispersion (20% vs. 0%), and the central age of the sample (~75 Ma vs. ~83 Ma), as well as modify the mixture model ages and the number of peaks identified during deconvolution [16–18].

Fission track analysis is a geochronological method that is highly dependent on choices made by an analyst. Therefore, evaluating, monitoring for, and mitigating fission-track analyst bias should not be marginalized by users of fission-track data [19, 20]. Complex Earth system interactions often produce complex datasets that are difficult to interpret or understand if we do not have all the necessary information at our disposal. Consequently, fission track datasets that are 'too good' with perfect 100% χ^2 passes should not be expected in every case and require scrutiny like any other heteroscedastic dataset. Imagine a hypothetical scenario where an old crystalline basement sample yields Precambrian apatite FT ages. Most of the time we would assume a priori that this sample *should* yield a single age population

and pass the χ^2 test, since all apatites are derived from the same rock. The analyst may struggle to find suitable grains for measurement due to high track density in many grains on the mount and select those considered easier to measure, unconsciously biasing towards younger ages. The analyst could also only count some, but not all, of the tracks in the counting area, which would yield low N_s totals and magnify the individual age errors. Ultimately, the central age would be more or less representative for the sample and we would have a χ^2 pass. The sole source apatite assumption and omission of apatite elemental data would implicitly validate our glowing statistics—but this may not be valid.

"Mistakes often come from assuming something is true just because there is little or no evidence against it."—R. Galbraith, *Statistics for Fission Track Analysis*, p. 190.

In addition to analyst bias, analytical canon has been to measure (up to) 20 age grains and 100 track lengths per sample. The χ^2 test is sensitive to sample size. The power of χ^2 increases when sample size is large enough and the absolute differences between individual ages become a progressively smaller proportion of the expected or 'true' value [21]. This means that any small deviations from an assumed 'true' age model in the dataset may appear statistically significant and provide support for the addition of mixture model age peaks [21]. Basically, the alternative to the null hypothesis is a very general one when n is large. Therefore, the higher the number of grains analyzed for an FT dataset, there is greater potential to fail the χ^2 test due to sample size [2], which is even more likely if single-grain ages are high precision. The reader is referred to [1, 2, 13, 19, 21] for further discussion of these topics.

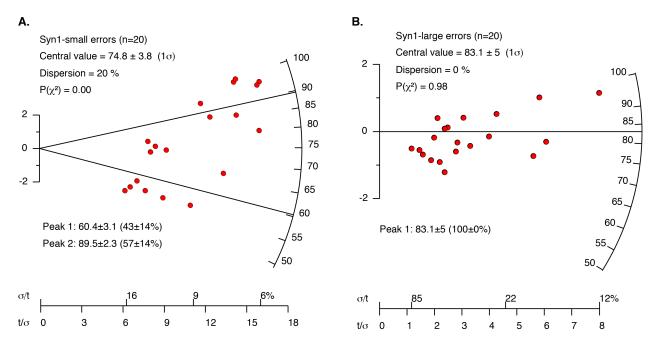


FIGURE 1. (A) Radial plot for synthetic sample with precise single-grain ages. (B) Radial plot for same synthetic sample as (A) with lower precision single-grain ages. Grain ages were held fixed whereas errors were randomly assigned for both samples. Plots created with DensityPlotter v. 8.4 [18].

— Thanks to Rex Galbraith and Nathan Cogné for comments and discussion

References

- [1] Galbraith RF (2005) Statistics for Fission Track Analysis (Chapman and Hall/CRC, Taylor and Francis Group, Boca Raton, FL) p 219.
- [2] McDannell KT, Issler DR, & O'Sullivan PB (2019) Radiation-enhanced fission track annealing revisited and consequences for apatite thermochronometry. *Geochimica et Cosmochimica Acta* 252:213-239.
- [3] Galbraith RF & Laslett GM (1993) Statistical models for mixed fission track ages. *Nuclear Tracks and Radiation Measurements* 21:459-470.
- [4] Carlson WD, Donelick RA, & Ketcham RA (1999) Variability of apatite fission-track annealing kinetics: I. Experimental results. *American mineralogist* 84(9):1213-1223.
- [5] O'Sullivan PB & Parrish RR (1995) The importance of apatite composition and single-grain ages when interpreting fission track data from plutonic rocks: a case study from the Coast Ranges, British Columbia. *Earth and Planetary Science Letters* 132:213-224.
- [6] Ravenhurst CE, Roden MK, Willett SD, & Miller DS (1993) Dependence of fission track annealing kinetics on apatite crystal chemistry. *Nuclear Tracks and Radiation Measurements* 21(4):622.
- [7] Carpéna J, Kienast J-R, Ouzegane K, & Jehanno C (1988) Evidence of the contrasted fission-track clock behavior of the apatites from In Ouzzal carbonatites (northwest Hoggar): The low-temperature thermal history of an Archean basement. *Geological Society of America Bulletin* 100:1237-1243.
- [8] Green P, Duddy I, Gleadow A, Tingate P, & Laslett G (1985) Fission-track annealing in apatite: track length measurements and the form of the Arrhenius plot. *Nuclear Tracks and Radiation Measurements* 10(3):323-328.
- [9] Green P, Duddy I, Gleadow A, Tingate P, & Laslett G (1986) Thermal annealing of fission tracks in apatite: 1. A qualitative description. *Chemical Geology: Isotope Geoscience section* 59:237-253.
- [10] Issler D, Grist A, & Stasiuk L (2005) Post-Early Devonian thermal constraints on hydrocarbon source rock maturation in the Keele Tectonic Zone, Tulita area, NWT, Canada, from multi-kinetic apatite fission track thermochronology, vitrinite reflectance and shale compaction. *Bulletin of Canadian Petroleum Geology* 53(4):405-431.
- [11] Issler DR, Lane LS, & O'Sullivan PB (2018) Characterisation, interpretation and modelling of multi-kinetic apatite fission track data using elemental data. (Geological Survey of Canada, Scientific Presentation 94, 1 sheet, https://doi.org/10.4095/311302).
- [12] Schneider DA & Issler DR (2019) Application of low-temperature thermochronology to hydrocarbon exploration. *Fission-Track Thermochronology and its Application to Geology*, eds Malusa MG & Fitzgerald P (Springer International Publishing, New York), 1 Ed, pp 315-333.
- [13] Vermeesch P (2017) Statistics for LA-ICP-MS based fission track dating. Chemical Geology 456:19-27.
- [14] Cogné N, Chew DM, Donelick RA, & Ansberque C (2020) LA-ICP-MS apatite fission track dating: A practical zeta-based approach. *Chemical Geology* 531.
- [15] Galbraith RF (1990) The radial plot: graphical assessment of spread in ages. *International Journal of Radiation Applications and Instrumentation. Part D. Nuclear Tracks and Radiation Measurements* 17(3):207-214.
- [16] Vermeesch P (2009) RadialPlotter: a Java application for for fission track, luminescence, and other radial plots. Radiation Measurements 44:409-410.
- [17] Galbraith R & Green P (1990) Estimating the component ages in a finite mixture. *International Journal of Radiation Applications and Instrumentation. Part D. Nuclear Tracks and Radiation Measurements* 17(3):197-206.
- [18] Vermeesch P (2012) On the visualisation of detrital age distributions, Chemical Geology 312:190-194.
- [19] Donelick RA, O'Sullivan PB, & Ketcham RA (2005) Apatite fission-track analysis. *Reviews in Mineralogy and Geochemistry* 58(1):49-94.
- [20] O'Sullivan P (2018) Statistical evaluation (and manipulation) of fission-track single-grain age data: an analyst's "unbiased" perspective. in *16th International Conference on Thermochronology* (Thermo2018, Palais Salfeldt, Quedlinburg, Germany).
- [21] Vermeesch P (2019) Statistics for Fission-Track Thermochronology. Fission-Track Thermochronology and its Application to Geology, eds Malusa MG & Fitzgerald P (Springer International Publishing, New York), 1 Ed, pp 109-122.

This technical note is a preprint and has not been submitted for journal publication. Please be aware that this manuscript was informally commented on but did not undergo formal peer review. These notes are meant for a general Earth science audience and users of thermochronology data. Please consult listed references for more detailed discussion of these topics.

Subsequent versions of this manuscript may have slightly different content.

Feel free to contact the author—all feedback is welcome