Paleoceanographic implications of submarine scours and contourites around ancient volcanoes, eastern Great Australian Bight

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ABSTRACT

Thermohaline oceanic currents control local and global variations in climate, vegetation, and biodiversity. Paleoceanographic studies typically use biostratigraphic and geochemical proxies to reconstruct the dynamics of these currents in Earth's ancient oceans. Here we use 2D seismic reflection data to interrogate the Middle Eocene-to-Recent stratigraphic record of ocean current evolution within the Ceduna Sub-basin, which is located in the eastern Great Australian Bight. These data show that 10–90 m deep, up to 3 km wide erosional scours, and <50 m thick, sediment wave-like bedforms, probably generated by contour currents, are developed throughout this carbonate-dominated succession. The scours are particularly well-developed at one specific stratigraphic level, and define 'moats' encircling Middle Eocene shield volcanoes, which, at that time, formed bathymetric highs. We suggest that sediment erosion, transport, and deposition may record Middle Eocene initiation of the Leeuwin Current, one of the most important ocean currents in the southern hemisphere. Deepest seabed scouring may reflect middle Miocene waxing of the so-called 'proto-Leeuwin Current', possibly driven by changes in ocean circulation patterns caused by the Miocene Global Optimum. The results of this seismic reflection-based study are consistent with results derived from other paleoceanographic proxies, thereby highlighting the key role seismic reflection data have in understanding the occurrence, geographical distribution, and significance of ancient ocean currents.

end of abstract

By influencing seawater temperature and salinity, ocean current activity controls regional and global trends in climate and biodiversity. Determining past changes in thermohaline circulation patterns in the world's oceans is thus important to understanding how climate and biodiversity varied in deep time and, therefore, may change in the future (cf. "geological analogues" of IPCC, 2007; see also e.g., Henderson, 2002; Wunsch, 2002; Rahmstorf, 2003; Wyrwoll et al., 2009). Paleooceanographic analysis commonly relies on biostratigraphic and geochemical proxy data, which: (i) are expensive to collect, and typically only collected on academic scientific cruises (e.g. IODP); (ii) are spatially limited (i.e. collected over a relatively small area within a discrete stratigraphic intervals); and (iii) do not typically provide a physical (i.e. stratigraphic) record of the initiation, extent, and decay of oceanographic currents (e.g. von Blackenburg, 1999; Henderson, 2002).

Seismic reflection data, which can provide relatively high-resolution images of the Earth's subsurface (e.g. Cartwright and Huuse, 2003), provide an additional but hitherto underutilised tool for determining the basin-scale dynamics of ancient ocean currents (e.g., Boldreel et al., 1998; Davies et al., 2001; Due et al., 2006; Hohbein et al., 2012). For example, seismic data can image large (up to 10's of metre thick, and several hundred-to-tens of kilometres long), contour current-driven bedforms, which document protracted transport and deposition of biogenic or detrital sediment (e.g. Rebesco and Stow, 2001; Stow et al., 2003), and/or scours of comparable depth, width, and length that record periods of net-seabed erosion and sediment redeposition. If integrated with other, more commonly used paleoceanographic proxies, seismic reflection data could form a key part of the oceanographer's toolkit.

In this study we use 2D seismic reflection data from the Ceduna Sub-basin, in the eastern Great Australian Bight to map seismic-scale, Middle Eocene-to-Recent scours and bedforms preserved adjacent to Middle Eocene volcanoes. We suggest these scours and bedforms reflect Middle Eocene initiation of the current now known as the Leeuwin Current. Although its age of initiation is debated, the present Leeuwin Current is the longest (5000 km) and one of the most important ocean currents in the southern hemisphere (see below). By providing physical evidence for ocean current activity, the results of our seismic reflection-based analysis at least broadly support studies proposing a Middle Eocene initiation age for the Leeuwin Current. More generically, our study opens up the exciting possibility that industry seismic reflection data, which are freely available for academic use along this and many other continental margins, if integrated with biostratigraphic and geochemical proxies, can help improve our understanding of deep-time dynamics of the Earth's ancient oceans.

Geological Setting. The Ceduna Sub-basin is located in the Bight Basin, offshore southern Australia (Fig. 1A), and formed in response to Jurassic-to-Early Cretaceous rifting and Early Cretaceous-to-Recent, post-rift thermal subsidence. Numerous submarine volcanoes were emplaced during earliest Middle Eocene magmatism at *ca*. 42 Ma (the 'Bight Basin Igneous Complex'; Schofield & Totterdell, 2008; Jackson, 2012; Magee et al., 2013), with the volcanoes onlapped by the fully marine, Middle Eocene-to-Recent, carbonate-dominated Nullarbor Limestone (Fig. 1B) (Schofield & Totterdell, 2008). Schofield & Totterdell (2008) and Jackson (2012) identify numerous 'scours' in the Nullarbor Limestone, although they do not explore their age or origin, their genetic relationship to the Middle Eocene volcanoes, or the potential paleoceanographic significance of their causal current. There are no direct constraints on Middle Eocene-to-Recent water depths in the distal Ceduna Sub-basin, although Jackson (2012) suggests the basin deepened to a few hundred metres during Middle Eocene flooding (see also McGowran et al., 1997; Shafik, 1983; 1990). The interpreted relatively deep-marine setting is consistent with water depth indicators provided by the heights of Eocene clinoforms preserved along the northern basin margin (Feary and James, 1995, 1998; McGowran et al., 1997). The Ceduna Sub-basin lies outboard of the modern day shelf edge in water depths of 200–4000 m; the seabed in the study area is below the influence of the modern Leeuwin Current, which extends to depths of 300 m (Feng et al., 2009).

Paleoceanography of the Great Australian Bight. The Leeuwin Current is the longest (5000 km) coastal current in the world and one of the most important ocean currents in the southern hemisphere (Fig. 1A). It is connected to and thus samples the global thermohaline system via the Indonesian Gateway (Feng et al., 2009), flowing southwards along the western coast of Australia and thereafter eastwards into the Great Australian Bight (Fig. 1A) (Cresswell and Golding, 1980). Transporting warm, low-salinity waters derived from the South Equatorial Current, the Leeuwin Current contributes to climatic variations and vegetation patterns (e.g., Caputi, 2001; Feng et al., 2009; Wyrwoll et al., 2009), and continent-scale biodiversity patterns by transporting otherwise low-latitude fauna to anomalously high latitudes (e.g., Cann and Clarke, 1993; McGowran et al., 1997; Passlow et al., 1997).

The Quaternary extent and dynamics of the Leeuwin Current in the Great Australian Bight are relatively well understood (Cresswell and Golding, 1980; Feng et al., 2009; Wyrwoll et al., 2009), whereas its timing of initiation, and its pre-Quaternary eastward 'reach' into the Great Australian Bight, remain uncertain. Based on their discovery of warm-water, Eocene microfauna in the Otway Basin, McGowran et al. (1997) suggested an even older, Middle Eocene age of initiation for the Leeuwin Current in the Great Australian Bight, arguing these fauna were likely derived from warm low-latitudes (i.e. via the proto-Leeuwin Current) and not cold high-latitudes (i.e. via the Flinders Current) (Fig. 1A); this interpretation is supported by fully coupled climate model simulations (Huber et al., 2004). However, pre-Early Oligocene initiation of the Leeuwin Current is disputed given that opening of the Tasmanian Gateway, which facilitated eastwards flow of warm ocean waters along southern Australia into the Pacific, did not occur until the Early Oligocene (e.g. Stickley et al., 2004; Wyrwoll et al., 2009). These studies together suggest the proto-Leeuwin Current in some way shaped the Eocene-to-Recent stratigraphic evolution of the Great Australian Bight shelfal regions, although the stratigraphic expression of time-equivalent, basin-centre units, which should also record the initiation and influence of this important ocean current, has yet to be documented.

Although Middle Eocene initiation of the Leeuwin Current is disputed, there is evidence from IODP Leg 182 that anomalously warm waters entered the Great Australian Bight during the middle Miocene, possibly in response changes in ocean circulation driven by the Miocene Climatic Optimum (e.g. Savin et al., 1975; Feary and James, 1995, 1998; Gourley and Gallagher, 2004). For example, McGowran et al. (1997) document the ‘Little Barrier Reef’, a thick (350 m), laterally extensive (>475 km long), rimmed carbonate platform developed along the northern margin of the Bight Basin (Fig. 1A).

Data and methods. Our dataset consists of 109 time-migrated, zero-phase, 2D seismic reflection lines that have a cumulative line length of *c*. 13,000 km and cover *c*. 44,000 km^2 of the central Ceduna Sub-basin (Fig. 1A). The NW- and NE-trending lines are spaced 4–16 km and 4–8 km, respectively (Fig. 2). The Gnarlyknots-1a borehole constrains the age of four key seismic reflections (Horizons A–D; Figs 1B and 3). This borehole also contains well-log data, indicating the Nullarbor Limestone has a p-wave velocity of 2100 m s^{-1}, thereby allowing us to convert measurements in milliseconds two-way time (ms TWT) to metres (cf. Espurt et al., 2009). Extrusive igneous bodies were identified and mapped using the geometric and geophysical criteria outlined by Totterdell & Schofield (2008), Jackson (2012) and Magee et al. (2013). Intra-Nullarbor scours are characterised by erosional surfaces that truncate and are onlapped by, underlying and overlying reflections respectively (Fig. 3). Although seismic data are only 2D and relatively widely spaced, scours and spatially related ‘hummocky’ seismic facies are typically imaged on and can be mapped between, several adjacent seismic lines. In particular, it is clear that the scours are only developed adjacent to and define ‘moats’ that encircle the volcanic vents (see below and Fig. 2).

Description of intra-Nullarbor seismic facies. We define two main stratal units within the Nullarbor Formation (SU1-2), separated by a major erosional surface (Horizon D; Figs 1 and 3). The base of SU1 is defined by a high-amplitude, laterally-continuous, positive seismic reflection defining the contact between the Pidinga Formation and the Nullarbor Limestone (i.e. Horizon B), or a series of volcanoes (i.e. Horizon C) (Figs 1B and 3). SU1 comprises the lower part of the Nullarbor Limestone and, away from the volcanoes, is typically characterised by low-to-moderate amplitude, parallel-to-sub-parallel, very continuous reflections (Fig. 3). Closer to the volcanoes (i.e. <2 km), a series of gently-dipping (<4°) reflections are developed in SU1, and these locally display bilateral downlap onto the underlying reflections and thus define convex-up, ‘mounded’ bodies (Figs 3B-D). These inclined reflections, which either dip towards (Fig. 3A) or away (Figs 3C–D) from the volcanoes, typically overlie low-angle (<6°) erosional surfaces, which are up to 2 km wide and display up to 100 m of relief; these surfaces pass laterally into (seismically) conformable surfaces (Fig. 3).

The base of SU2 is locally defined by a major erosion surface along which numerous scours are developed (Horizon D; Figs 1–3). These scours locally define a series of ‘moat-like’ features that fully or partly encircle 15 of the 57 vents present in the Ceduna Sub-basin (e.g. Fig. 2). The scours have a relief of 10–90 m, extend <3 km from the volcanoes, and their flanks dip 0.1–6.1°, being best-developed around volcanoes that are typically >200 m tall. Some of the scours are asymmetric, consisting of a long, gently dipping outer margin inclined towards the vents and a shorter, more steeply-dipping surface that dips away from the vents (Figs 3B–D). Our 2D seismic data do not allow us to confidently determine if the scours are preferentially developed on one side of the vents (Fig. 2). Two main types of seismic facies fill the scours: (i) high-amplitude, ‘hummocky’ reflections, which have a relief of up to 50 m and a distance of 100–200 m between adjacent hummock crests (Fig. 3B); and (ii) low-to-high amplitude, gently-dipping (<2°), moderately discontinuous to laterally-continuous reflections (Fig. 3). The upper part of SU2 is dominated by low-to-high amplitude, flat-lying to gently-dipping (<2°), laterally-continuous reflections (Fig. 3), with erosionally based packages of chaotic reflections locally being developed.

Interpretation of intra-Nullarbor features. Based on their development in a fully marine succession and given that post-Middle Eocene water depths were probably at least several hundred metres (see Jackson, 2012), it is unlikely the intra-Nullarbor scours and hummocks formed subaerially. Furthermore, the coeval basin margin, which was likely located several hundred kilometres to the north, was carbonate-dominated and constructional (Fig. 1; see also Feary and James, 1995, 1998), suggesting little or no sediment bypass was occurring, and that voluminous, strongly erosional, deep-water gravity currents, such as turbidity currents, did not form the intra-Nullarbor scours and hummocks.

Based on their development in a fully marine succession deposited in several hundreds of metres of water, our preferred interpretation is that the scours formed in response to ocean current-related incision of the seabed (Fig. 4). Such scours are common in modern seas and oceans, typically in association with channel-related contourite drifts (*sensu* Stow et al., 2002). The spatial restriction of Ceduna Sub-basin scours to within c. 3 km of the volcanoes, strongly suggests the volcanic edifices formed syn-incision bathymetric highs that perturbed the velocity structure of the causal ocean currents. We suggest this perturbation increased current turbulence and, most critically, seabed shear stress, driving localised erosion of the seabed immediately adjacent to the volcanoes (Fig. 4) (e.g. O'Reilly et al., 2003; MacLachlan et al., 2008). Submarine scours of broadly similar geometry, dimension, and origin are observed adjacent to igneous rock-cored bathymetric highs in the Pisces Reef system, Irish Sea, UK (Callaway et al., 2009), and in the Capel and Faust basins, offshore eastern Australia (Rollet et al., 2012).

We interpret that inclined and hummocky reflections developed throughout the Nullarbor Limestone represent dip-oblique and dip-parallel sections, respectively, through sediment wave- or contourite drift-like deposits (cf. Rebesco and Stow, 2001; Stow et al., 2003; Hohbein et al., 2012). Sedimentary bodies like this are commonly associated with seabed scours, being deposited when bottom current energy is low enough to permit sediment reworking within bedforms (Fig. 4) (e.g. Stow et al., 2002). Like the scours, intra-Nullarbor bedforms are spatially restricted to within a few kilometres of the vents, suggesting they too formed due to volcano-driven perturbations in ocean current velocity and seabed shear stress. In their case, however, an increase in seabed shear stress was only sufficient to rework sediment and not deeply erode the seabed (Fig. 4).

Implications for the paleoceanographic development of the Great Australian Bight. Although the Gnarlyknots-1a borehole penetrates the Nullarbor Formation, no biostratigraphic data were collected; as a result, we cannot constrain the age of intra-Nullarbor scours, associated strata, or indeed, the causal current more tightly than 'Middle Eocene-to-Recent'.

Furthermore, no paleobathymetric data (e.g. benthic foraminifera) were collected in Gnarlyknots-1a, meaning we have no direct constraints on water depth variations during the Middle Eocene-to-Recent, and thus the depth of formation of the ocean current-related scours and bedforms remains uncertain (see also discussion in Jackson, 2012). Our relatively widely spaced 2D seismic reflection data also do not allow us to confidently determine if intra-Nullarbor scours are preferentially developed on one, most likely the down-current (i.e. lee) side of the seabed obstruction (i.e. the volcanoes), or if the associated bedforms are best-developed on one, most likely the up-current side of the vents, and display down-current accretion (e.g. Callaway et al., 2009; Rollet et al., 2012). Because of this, we do not know the dominant direction of the causal current.

Notwithstanding these limitations, it is informative to discuss the implications of our study for the Paleogene paleoceanographic evolution of the eastern Great Australian Bight. Using biostratigraphic proxy data from the Bight and Otway basins, McGowran et al. (1997) suggest the initiation of eastwards protrusion of a so-called ‘proto-Leeuwin Current’ into the Great Australian Bight during the Middle Eocene, with further evidence for its presence in the middle Miocene (Fig. 1A; see also Feary and James, 1995, 1998). This interpretation was, however, challenged by Wyroll et al. (2009), who suggest these fauna may simply record locally elevated sea surface temperatures unrelated to the initiation of a plate-scale thermohaline current. We here suggest Middle Eocene-to-Recent scours and bedforms developed in the Ceduna Sub-basin are associated with Paleogene initiation of an oceanographic current, which we link to the postulated proto-Leeuwin Current (Fig. 4). More specifically, we propose these features record late Middle Eocene initiation and subsequent fluctuations in the strength and erosivity of, the current. The major intra-Nullarbor erosion surface (Horizon D), for example, which is best-developed immediately adjacent to the volcanoes, may represent intensification or ‘waxing’ of the proto-Leeuwin Current (Fig. 4). Despite a lack of age data, we tentatively suggest this erosional event, and related bedforms, may be the stratigraphic expression of the Miocene Climatic Optimum-related event proposed by Feary and James (1995, 1998), during which time anomalously warm waters encroached eastwards into the Great Australian Bight from western Australian (see also Savin et al., 1975; Gourley and Gallagher, 2004).

Our study shows that seismic reflection data can image erosional and depositional features that provide a physical stratigraphic record of ancient ocean currents. We demonstrate that by placing these features into a broad chronostratigraphic framework, we can complement rather sparse micro-faunal evidence, and gain important insights into the timing of onset of major ocean currents. Data limitations notwithstanding, our seismic reflection-based approach does, at the very least, provide a clear hypothesis testable with future scientific drilling (e.g. IODP). Seismic reflection data allow erection of a physical, stratigraphic framework and may be an essential part of the paleoceanographer's toolkit. Future work should focus on detailed mapping of seismic reflection datasets from, for example, the western Bight Basin and Otway Basin; this may reveal similar, age-equivalent current-formed stratigraphic features, thus raising the possibility that the proto-Leeuwin Current extended further eastwards and influenced faunal distribution and potentially climate over a wider area than currently assumed.

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FIGURE CAPTIONS

Fig 1. (A) Map illustrating the geographical setting of the study area. The area covered by 2D seismic reflection data is outlined by a solid black line. Inset shows the key modern oceanographic currents developed along the western and southern Australian margin. LC=Leeuwin Current; ACC=Antarctic Circumpolar Current; WAC=West Australia Current; FC=Flinders Current. OB=Otway Basin. Modified from Jackson (2012); oceanographic currents from Bilj et al. (2013). (B) Simplified stratigraphic column based on data from boreholes Gnarlyknots-1A and Potoroo-1. Key seismic horizons (A–D) and seismic units (SU1-2) are indicated. The stratigraphic occurrence of intrusion and extrusive components of the Bight Basin Igneous Complex (BBIC) are shown. Modified from Jackson (2012).

Fig. 2. Time-structure map of the Horizon D merged with a time-structure map of Horizon C; this illustrates the distribution of volcano summits (labelled 'v' and encircled by a solid white line) that rise above Horizon D, and intra-SU2 moats (labelled 'm'). sw=sediment wave crests. (A) and (B) are from the southern and northern parts of the study area respectively. See Figure 1A for location of map. Light grey lines indicate seismic reflection profiles.

Fig. 3. (A)–(D) Interpreted seismic profiles illustrating the geometry, scale and relationship between extrusive volcanic features of the Bight Basin Igneous Complexes and intra-Nullabor Limestone scours (moats) and bedforms. Locations of the seismic lines are shown in Figure 2. Uninterpreted versions of sections are available in DRI1.

Fig. 4. (A–C) Schematic diagrams illustrating the evolution of intra-Nullarbor Limestone scours in the Great Australia Bight. The vertical dashed line in the 'current strength' column indicates the threshold for sediment erosion/non-deposition; these conditions occur at the onset of T2.

Data Repository Item (DR1). Uninterpreted versions of the seismic profiles presented in Fig. 3.

Fig. 1

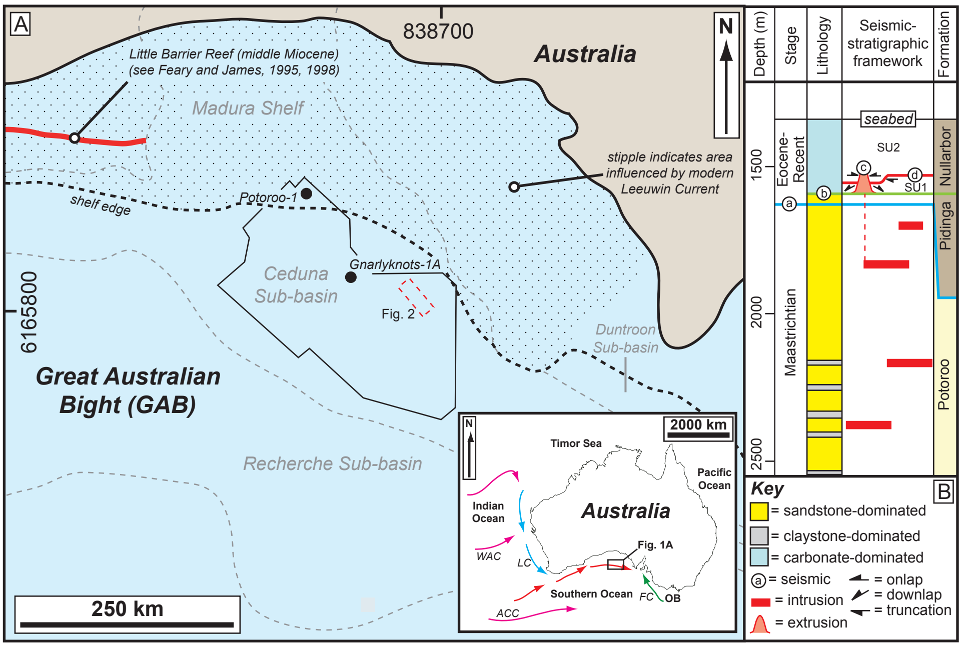
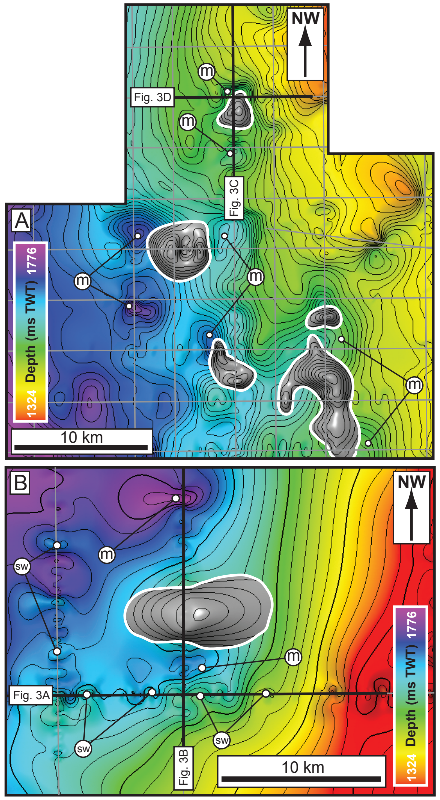
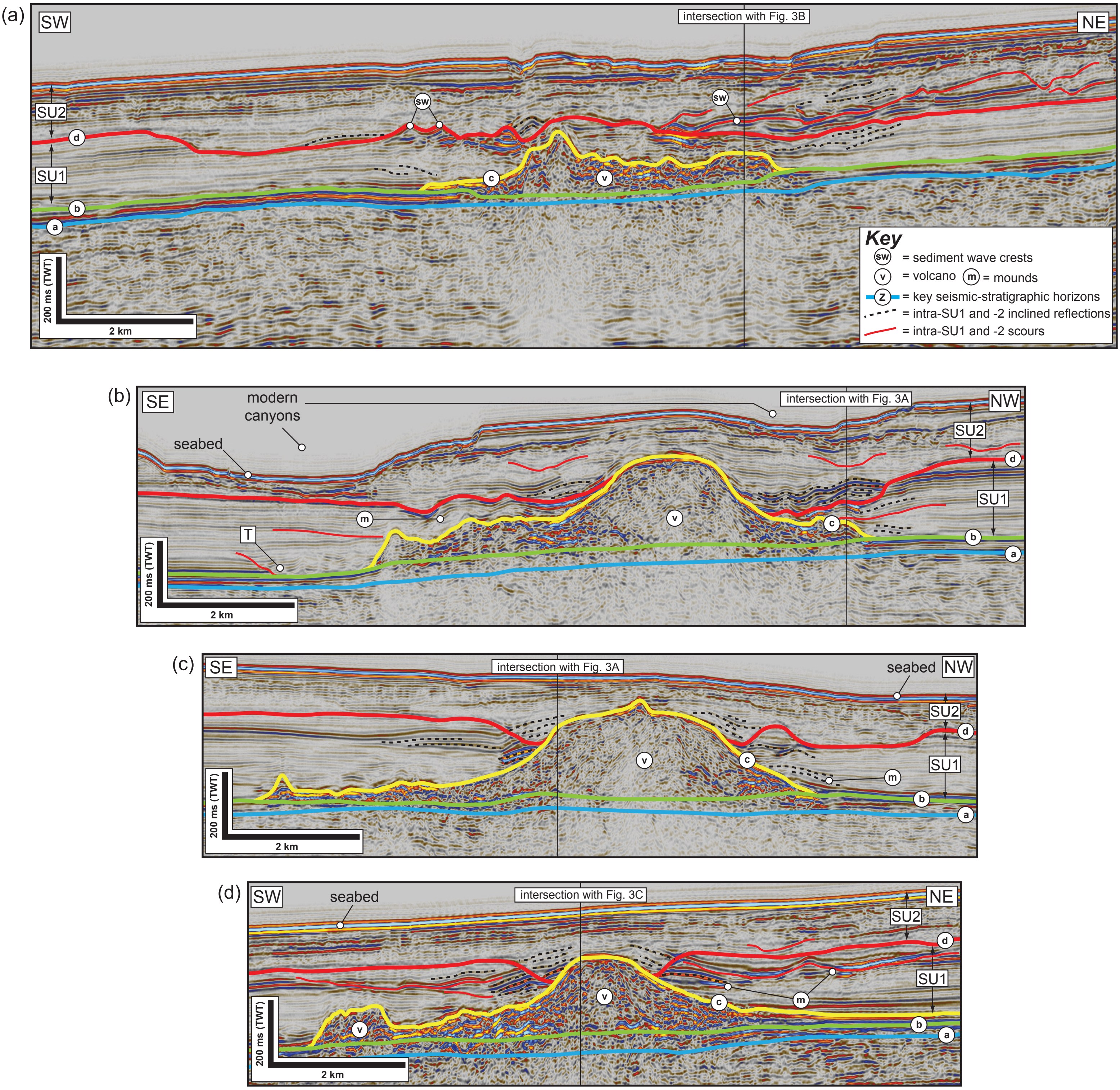

A
838700
Australia
N
Little Barrier Reef (middle Miocene)
(see Feary and James, 1995, 1998)
Madura Shelf
stipple indicates area influenced by modern Leeuwin Current
shelf edge
Potoroo-1
Gnarlyknots-1A
Ceduna Sub-basin
Fig. 2
Duntroon Sub-basin
6165800
Great Australian Bight (GAB)
Recherche Sub-basin
250 km
2000 km
Timor Sea
Pacific Ocean
Indian Ocean
Australia
WAC
LC
Fig. 1A
Southern Ocean
FC
OB
ACC
Depth (m)
Stage
Lithology
Seismic-stratigraphic framework
Formation
seabed
SU2
SU1
Eocene-Recent
Maastrichtian
Nullarbor
Pidinga
Potoroo
1500
2000
2500
B
Key
= sandstone-dominated
= claystone-dominated
= carbonate-dominated
= seismic
= intrusion
= extrusion
= onlap
= downlap
= truncation

Fig. 2



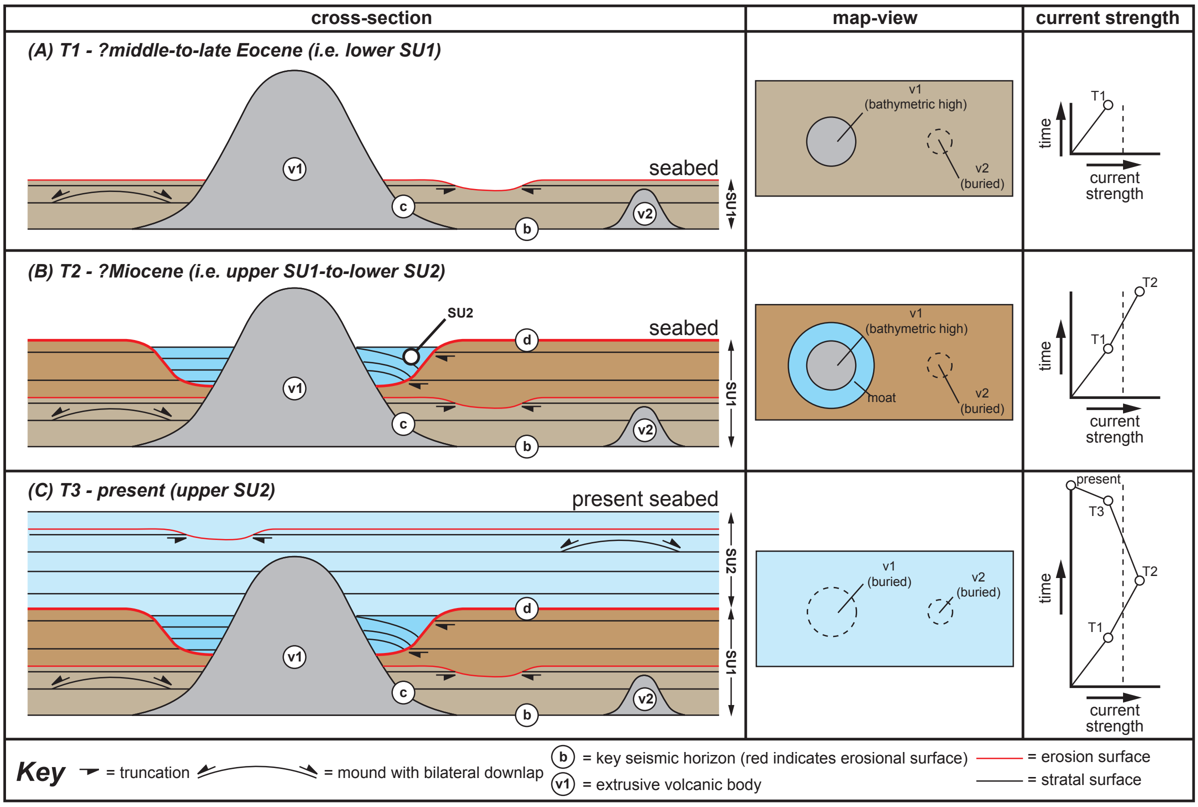
A
NW
Fig. 3D
Fig. 3C
m
1324 Depth (ms TWT) 1776
10 km
B
NW
sw
m
Fig. 3A
Fig. 3B
1324 Depth (ms TWT) 1776
10 km

Fig. 3



(a)
SW
NE
intersection with Fig. 3B
SU2
SU1
sw
c
v
d
b
a
200 ms (TWT)
2 km
Key
= sediment wave crests
= volcano
= mounds
= key seismic-stratigraphic horizons
= intra-SU1 and -2 inclined reflections
= intra-SU1 and -2 scours
(b)
SE
NW
modern canyons
seabed
intersection with Fig. 3A
SU2
SU1
m
T
(c)
intersection with Fig. 3A
seabed
(d)
intersection with Fig. 3C

Fig. 4


cross-section
map-view
current strength
(A) T1 - ?middle-to-late Eocene (i.e. lower SU1)
seabed
v1
v2
c
b
SU1
v1 (bathymetric high)
v2 (buried)
T1
time
current strength
(B) T2 - ?Miocene (i.e. upper SU1-to-lower SU2)
SU2
seabed
d
moat
T2
(C) T3 - present (upper SU2)
present seabed
v1 (buried)
v2 (buried)
present
T3
Key
= truncation
= mound with bilateral downlap
= key seismic horizon (red indicates erosional surface)
= extrusive volcanic body
= erosion surface
= stratal surface

Data Repository Item 1 (DR1)