

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25

Modeling transpiration with sun-induced chlorophyll fluorescence via water use efficiency and stomatal conductance

Huaize Feng¹, Tongren Xu¹, Xinlei He¹, Jingxue Zhao¹, Shaomin Liu¹

¹ State Key Laboratory of Earth Surface Processes and Resource Ecology, School of Natural Resources, Faculty of Geographical Science, Beijing Normal University, Beijing 100875, China

Key Points:

- SIF can be used to model transpiration via water use efficiency and stomatal conductance
- Mechanism models of SIF and transpiration outperformed linear model
- Air dryness has an important impact on the relationship between transpiration and SIF.

Abstract

Successfully applied in the carbon research area, sun-induced chlorophyll fluorescence (SIF) has raised the interest of researchers from the water research domain. However, the mechanism between SIF emitted by plants and transpiration (T) has not been fully explored. To improve the understanding of the relationship between SIF and T, we developed two SIF-T models, the WUE model and the conductance model, based on carbon-water coupling framework. Hourly data were used to build and validate the models we developed. Correlation analysis shows that the T modeled by our model outperforms the traditional empirical linear model with higher R^2 and lower RMSE. The models we built further show the potential and mechanism in estimating water flux by SIF.

1 Introduction

Evapotranspiration (ET) is not only a pipeline of the water cycle in the air but also an important influence factor of energy balance as a carrier of latent heat. Previous works indicated transpiration (T) occupies a dominant position in evapotranspiration [Good *et al.*, 2015; Jasechko *et al.*, 2013]. In some ecosystems, T could reach 95% of the total ET [Stoy *et al.*, 2019]. T is also closely coupled with the carbon plant productivity [Kool *et al.*, 2014]. Therefore, an accurate understanding of the

26 spatiotemporal changes of T is crucial for understanding the substance and energy interactions
27 between the land surface and the atmosphere.

28 In this century, Sun-induced chlorophyll fluorescence (SIF) renewed the gross primary
29 production (GPP) estimation from ground to space [*Frankenberg and Berry, 2018; Ryu et al., 2019;*
30 *Schimel et al., 2019*]. Considering the connection between photosynthesis and transpiration, SIF
31 may serve as a pertinent constrain estimates for transpiration. [*Alemohammad et al., 2017; Jonard*
32 *et al., 2020*]. Recently, Empirical analysis based on ground and remote sensing SIF observation
33 showed SIF is strongly related to T, [*Lu et al., 2018*] reconstructed the full band SIF and exploit
34 the capacity of individual SIF bands and their combinations for deriving T with empirical linear
35 regression and Gaussian Process Regression model at Harvard Forest. [*Pagán et al., 2019*] used
36 radiation corrected GOME-2 SIF observations to diagnose transpiration efficiency understood as
37 the ratio between transpiration and potential evaporation worldwide. [*Maes et al., 2020*]
38 investigated the empirical link between SIF and T using satellite SIF (GOME-2 and OCO-2) and
39 SCOPE model at sites from FLUXNET. However, the studies mentioned above relate T with SIF
40 empirically, not mechanistically. SIF is the light signal from the excited chlorophyll a molecules
41 after absorption of photosynthetically active radiation. The information about the electron
42 transport (J) from photosystem II to photosystem I contained in SIF makes the signal a powerful
43 tool to predict GPP [*Gu et al., 2019; Köhler et al., 2018; Zhang et al., 2014*]. Furthermore, the
44 essential of understanding the SIF-T relationship should lie in the coupling between the carbon
45 and water cycles.

46 The carbon and water cycles between the biosphere and atmosphere are strongly coupled
47 [*Gentine et al., 2019*]. The trade-off between photosynthesis and water vapor loss is arguably the
48 most central constraint on plant function [*Wolz et al., 2017*]. Water-use efficiency (WUE) and

49 stomatal conductance (g_s , or canopy conductance, g_c) are two key metrics of carbon - water
50 coupling. WUE is defined as the amount of carbon assimilated relative to water use [Leakey *et al.*,
51 2019]. [Maes *et al.*, 2020] reported WUE and related variables show a most important impact on
52 the SIF-T relation. Plants take in carbon dioxide and breath out water through stomata
53 simultaneously. Stomata played a key role in the carbon-water coupling, even the whole Earth
54 System [Berry *et al.*, 2010]. By analyzing the empirical link of SIF and canopy conductance, [Shan
55 *et al.*, 2019] reported the empirically linear linkage between g_s and SIF data from C3 forest,
56 cropland, and grassland ecosystems, and T calculated by SIF-based conductance agreed well with
57 ET observed by flux towers. However, it is still unclear how can WUE and g_s be used to model T
58 by SIF mechanistically.

59 In this paper, two carbon-water coupling indicators: water use efficient and stomatal
60 conductance are introduced to clarify the physical relevance between SIF and T. Using the concept
61 of these two indicators, two mechanism SIF-T models are built and tested based on hourly ground
62 observation from four sites, including two C4 and two C3 sites. The results are expected to improve
63 our understanding of the link between SIF and T.

64 **2 Materials and Methods**

65 2.1 Materials

66 SIF and corresponding observation (eg. weather variables, flux data, and vegetation indexes)
67 are acquired at four sites including two maize field sites (Daman from Heihe river basin, China,
68 DM; Huailai from Haihe river basin, China, HL), one temperate deciduous forest site (Harvard
69 Forest from AmeriFlux network, US, HF) and a subalpine conifer forest (Niwot Ridge from
70 AmeriFlux network, US, NR). The characteristics of these sites are summarized in Table 3. The
71 SIF measurements of DM and HL sites (760 nm) were conducted using a tower-based automatic

72 measurement system named “SIFSpec” [Du *et al.*, 2018] and retrieved using the 3FLD method
73 [Liu *et al.*, 2020]. The SIF data (745- to 758-nm) of NR were from a scanning spectrometer
74 (PhotoSpec) on the top of the 26 meters tower. SIF data in winter were abandoned in this study
75 because the subalpine trees at the NR site undergone significant physiological stress during the
76 cold climate [Magney *et al.*, 2019]. The SIF data of HF were retrieved from FluoSpec deployed
77 about 5 m above the canopy on the top of a tower and extracted by spectral fitting methods at 760
78 nm [Yang *et al.*, 2015].

79 Leaf area index (LAI) of all four sites was acquired from the MCD15A3H dataset with 4-day
80 and 500 m temporal-spatial resolution [Myneni *et al.*, 2015] and interpolated to hourly scale on
81 Google Earth Engine [Gorelick *et al.*, 2017]. Gross primary production (GPP) was separated from
82 net ecosystem exchange (NEE) following [Reichstein *et al.*, 2005] and [Lasslop *et al.*, 2010] via
83 the REddyProcWeb online tool ([https://www.bgcjena.mpg.de/bgi/index.php/Services](https://www.bgcjena.mpg.de/bgi/index.php/Services/REddyProcWeb)
84 [/REddyProcWeb](https://www.bgcjena.mpg.de/bgi/index.php/Services/REddyProcWeb)). One hour before the rainfall and six hours after the rainfall data were excluded
85 to minimize the influence of canopy interception.

86 T modeled by SIF is evaluated by T_{Zhou} partitioned from ET. [Zhou *et al.*, 2014] proposed an
87 index called underlying water-use efficiency (uWUE) by combining the optimal stomatal behavior
88 model with Fick's law. At the leaf scale, uWUE is defined as:

$$89 \quad uWUE_1 = \frac{GPP \times \sqrt{VPD}}{T} \quad (1)$$

90 Where VPD is the vapor pressure deficit, which is calculated from air temperature and relative
91 humidity of air. At the ecosystem scale, uWUE is written as:

$$92 \quad uWUE_a = \frac{GPP \times \sqrt{VPD}}{ET} \quad (2)$$

93 When water flux from the surface is fully occupied by transpiration from plants, $uWUE_a$ reaches
 94 its potential value $uWUE_p$. The $uWUE_p$ is assumed to be equaling to $uWUE_l$ and can be captured
 95 by the 95-quantile regression analysis of $uWUE_a$. Then T can be derived by:

$$96 \quad T_{Zhou} = \frac{GPP \times \sqrt{VPD}}{uWUE_p} \quad (3)$$

97 This method was developed based on flux tower data of 14 sites, including the HF site, and
 98 successfully applied to Heihe River Basin [Zhou *et al.*, 2018], including the DM site. Especially,
 99 at the DM site, T/ET estimated by the uWUE method agreed with the isotope method well during
 100 the peak growing season [Bai *et al.*, 2019].

101 2.2 WUE Model

102 SIF can be represented in the form of light use efficiency (LUE) model:

$$103 \quad SIF = APAR \Phi_F \Omega_c \quad (4)$$

104 where APAR stands the photosynthetically active radiation absorbed by photosynthetic
 105 pigments, Φ_F is the fluorescence quantum yield and Ω_c is the probability of SIF photon escaping
 106 from the canopy. Combined Eqn 4 with the LUE model of GPP (Eqn 5). We can obtain a linear
 107 model between SIF and GPP (Eqn 3):

$$108 \quad GPP = APAR LUE \quad (5)$$

$$109 \quad GPP = SIF \frac{LUE}{\Phi_F \Omega_c} \quad (6)$$

110 where LUE is the light-use-efficiency. A series of papers have reported the linear relationship
 111 between SIF and GPP by comparing GPP with satellite remote sensing SIF [Guanter *et al.*, 2014;
 112 Li *et al.*, 2018; Sun *et al.*, 2017] and field-based SIF observation [Liu *et al.*, 2017; Magney *et al.*,
 113 2019; Yang *et al.*, 2015]. Based on these works, the factor $\frac{LUE}{\Phi_F \times \Omega_c}$ can be set as a constant for a
 114 specific plant type, and GPP can be calculated by SIF directly. For C3 and C4 plants, WUE is

115 relatively stable. Many works derived T from GPP by treating WUE as a constant value during
 116 certain period [Scott and Biederman, 2017; Yang et al., 2015]. Under this assumption T can be
 117 calculated by SIF via the following equation:

$$118 \quad GPP_{\text{linear}} = k1 \text{ SIF} \quad (7)$$

$$119 \quad T_{\text{linear}} = k2 \text{ GPP}_{\text{linear}} \quad (8)$$

120 where k1, k2 are two parameters denoting $\frac{\text{LUE}}{\Phi_f \times \Omega_c}$ and WUE respectively. Eqn 8 is the theory base
 121 of the empirical linear relationship between SIF and T.

122 However, WUE is strongly affected by the dryness of air from leaf to ecosystem scale. The
 123 relationship between GPP and T improved significantly by incorporating the effects of VPD from
 124 diurnal to annual time scales [Beer et al., 2009; Zhou et al., 2014]. Moreover, [Jonard et al., 2020]
 125 pointed out the atmospheric demand for water helps explaining a lot variability in the SIF–T
 126 relationship at the ecosystem scale. Here we proposed a WUE model:

$$127 \quad GPP_{\text{WUE}} = k3 \text{ SIF} \quad (9)$$

$$128 \quad T_{\text{WUE}} = k4 \text{ VPD}^{k5} \text{ GPP}_{\text{WUE}} \quad (10)$$

129 where k3 is a parameter like k1. k4 is a parameter concluding information on water-use efficiency.
 130 k5 quantifies of the non-linear effect of VPD on k4 [Lin et al., 2018]. In this work, the VPD of air
 131 is used to describe the aridity stress at the canopy scale.

132 2.3 Conductance Model

133 Though the linear SIF-GPP relationship looks simple, the mechanism of the parameter LUE
 134 in Eqn 3 is complex. Φ_f and Ω_c are relatively stable value [Guanter et al., 2014; Tol et al., 2014],
 135 and LUE is often calculated by reducing potential LUE using several environmental factors, such
 136 as temperature, soil moisture [Yuan et al., 2007]. Previous work had also reported the hyperbolic
 137 relationship between SIF and GPP [Damm et al., 2015; Zhang et al., 2020]. Therefore, the

138 relationship between SIF and GPP is far more complicated than linear. As mentioned above, the
 139 link between SIF and GPP is because of the close relationship between SIF and electron transport
 140 rate (J). Benefit from the carbon-pump mechanism, the GPP of C4 plants is linearly related to J
 141 [Gu *et al.*, 2019]. For C4 plants, GPP can be derived by:

$$142 \quad J = a q_L \text{ SIF} \quad (11)$$

$$143 \quad \text{GPP}_{\text{gs}} = J/4 = \frac{a q_L \text{ SIF}}{4\Omega_c} \quad (12)$$

144 a is an empirical factor supposed to be a constant under ideal environments, q_L is the fraction of
 145 open Photosynthesis II reaction centers, indicating the 'traffic jam' in the electron transport
 146 pathway from Photosystem II to Photosystem I. Note that, the Eqn 11 is designed for broadband
 147 SIF for PSII. In this paper, single NIR band SIF is used instead by assuming a linear relationship
 148 between single-band SIF and full band SIF. q_L ranges from 0-1 and decreases with increased PAR
 149 [Baker, 2008; Gu *et al.*, 2019]. In this paper, q_L is derived by:

$$150 \quad q_L = \exp(-\beta \text{PAR}) \quad (13)$$

151 β is a parameter denoting the sensitivity of q_L to the illumination. Due to data restrictions, Ω_c for
 152 the near-infrared band SIF was set as a constant in our study. g_s for C4 plants is derived by inserting
 153 Eqn into the famous Ball-Berry model [Ball *et al.*, 1987]:

$$154 \quad g_s = m \frac{a q_L \text{ SIF}}{4\Omega_c} \text{Rh}/C_a + g_0 \quad (14)$$

155 Where m is an empirical slope parameter, which is often treated as a constant for a specific
 156 ecosystem [Miner *et al.*, 2017]. C_a is the ambient carbon dioxide concentration and g_0 is the
 157 minimum conductance which is set as 0. Lack of an efficient mechanism gathering CO_2 from the
 158 air, C3 plants much more rely on the stomata to absorb CO_2 for the Calvin cycle. For C3 plants,
 159 the relationship between SIF and GPP is also affected by the dark reactions [Gu *et al.*, 2019]:

$$160 \quad GPP_{gs} = a \frac{C_i - \Gamma^*}{4C_i + 8\Gamma^*} q_L \text{SIF} \frac{1}{\Omega_c} \quad (15)$$

161 C_i is the intercellular CO_2 concentration. Γ^* is the CO_2 compensation point in the absence of
 162 mitochondrial respiration, which can be set as a constant for specific plant type or calculated by
 163 air temperature [Katul *et al.*, 2010]. C_i in Eqn 10 can be eliminated by combining with Eqn 7 and
 164 Fick's law $GPP = gs \times (C_i - C_a)$, then we have:

$$165 \quad GPP_{gs} = a \frac{GPP/gs + C_a - \Gamma^*}{4(GPP/gs + C_a) + 8\Gamma^*} q_L \text{SIF} \frac{1}{\Omega_c} \quad (16)$$

166 gs and GPP_{gs} can be solved under the constraining of the optimality theory of stomatal behavior.
 167 [Cowan and GD, 1977; Katul *et al.*, 2010; Way *et al.*, 2014]. According to this theory, plants tend
 168 to adapt stomata to minimize the cost of water while maximizing carbon assimilation:

$$169 \quad f(gs) = GPP - \lambda ET \approx GPP - 1.6\lambda gs \text{VPD}/P \quad (17)$$

$$170 \quad \delta f(gs)/\delta(gs) = 0 \quad (18)$$

171 λ represents the marginal water cost of carbon assimilation. P is the air pressure. If we incorporate
 172 Eqn 16, 17, and 18, gs can be expressed as the function \mathcal{F} of SIF, q_L , λ , Γ , VPD and C_a :

$$173 \quad gs = \mathcal{F}(\text{SIF}, q_L, \lambda, \Gamma, \text{VPD}, C_a)$$

$$174 \quad = -\frac{a \text{SIF} q_L (4\Gamma - C_a)}{4(2\Gamma + C_a)^2} + \frac{a \text{SIF} q_L (2\Gamma + C_a - 3.2\lambda \text{VPD}) \sqrt{3.2\lambda \text{VPD} \Gamma (C_a - \Gamma) (2\Gamma + C_a - 1.6\lambda \text{VPD})}}{6.4\lambda \text{VPD} (2\Gamma + C_a)^2 (2\Gamma + C_a - 1.6\lambda \text{VPD})} \quad (19)$$

175 Finally, with SIF-based gs , T can be calculated by the two-source Penman-Monteith method
 176 [Leuning *et al.*, 2008]:

$$177 \quad Ac = Rn \times [1 - \exp(-0.5LAI) / \cos(SZA)] \quad (20)$$

$$178 \quad T_{gs} = \frac{\Delta Ac + \rho C_p \text{VPD} ga}{\Delta + \gamma \left(1 + \frac{ga}{gs}\right)} \quad (21)$$

179 Eqn 20 is the simple one-dimensional Beer's law model. Ac is the available energy of the canopy
 180 layer, Rn is net radiation, LAI is the leaf area index and SZA is the sun zenith angle. For Eqn 21,

181 where Δ is the rate of change of vapor pressure with temperature, γ is the psychrometric constant,
182 C_p is the specific heat of air; ρ is the density of liquid water and g_a is aerodynamic conductance.

183 2.3 Model Calibration

184 Parameters of both the WUE model and the conductance model need to be calibrated. GPP
185 and the water balance framework are used here to constrain the models. Total ET observed by
186 eddy covariance measurement is composed of plant transpiration (T) and soil evaporation (LEs):

$$187 \quad \quad \quad LE = T + LE_s \quad \quad \quad (22)$$

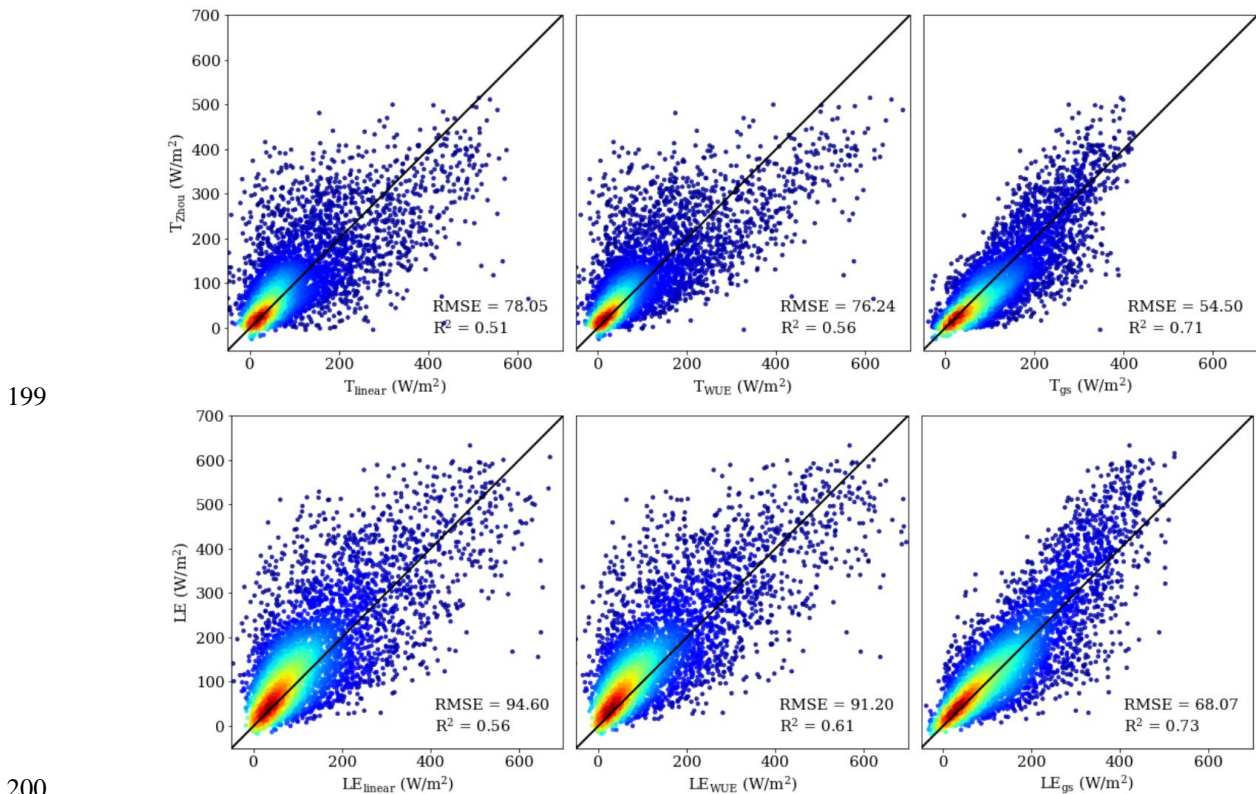
188 Considering the nonlinearity and complicity of the models, the shuffled complex evolution (SUE-
189 UA) algorithm [*Duan et al.*, 1994] is employed to fit parameters by maximizing the cost function
190 G:

$$191 \quad \quad \quad G = 0.4NSE(GPP_{ob}, GPP_{model}) + 0.6NSE(LE_{ob}, LE_{model}) \quad \quad \quad (23)$$

192 where NSE is the Nash-Sutcliffe efficiency coefficient. The subscript _{ob} means observed values
193 and variables with a subscript _{model} mean values are derived by models described above. All model
194 estimations and statistical analyses were performed with Python 3.8.3 [*Herman and Usher*, 2017;
195 *Houska et al.*, 2015]. The description and calculation of all variables mentioned above are listed
196 in Table 2.

197 **3 Results**

198 **3.1 Performance of SIF- T model**



200 **Figure 1.** Relationship of T modeled by SIF and reference latent heat (Hourly). Colors indicate
 201 the density of points (from sparse to dense: blue to red). R² denotes the coefficient of determination.
 202 The unit of root-mean-square deviation (RMSE) is W/m².
 203

204 Scatter plots between reference T_{Zhou} and T modeled by SIF are shown in Fig1 (upper row).

205 In general, both the WUE method and the conductance method improve the ability of SIF in
 206 modeling T. Compared with simple linear regression with R² = 0.51 and RMSE = 78.05 W/m², the
 207 WUE model has R² = 0.56 and RMSE = 76.24 W/m², while the conductance model has the highest
 208 R² = 0.71 and much lower RMSE = 54.50 W/m². The linear model and the WUE model tend to
 209 overestimate T at the high values area, while most points of T_{gs} fall near the 1:1 line.
 210

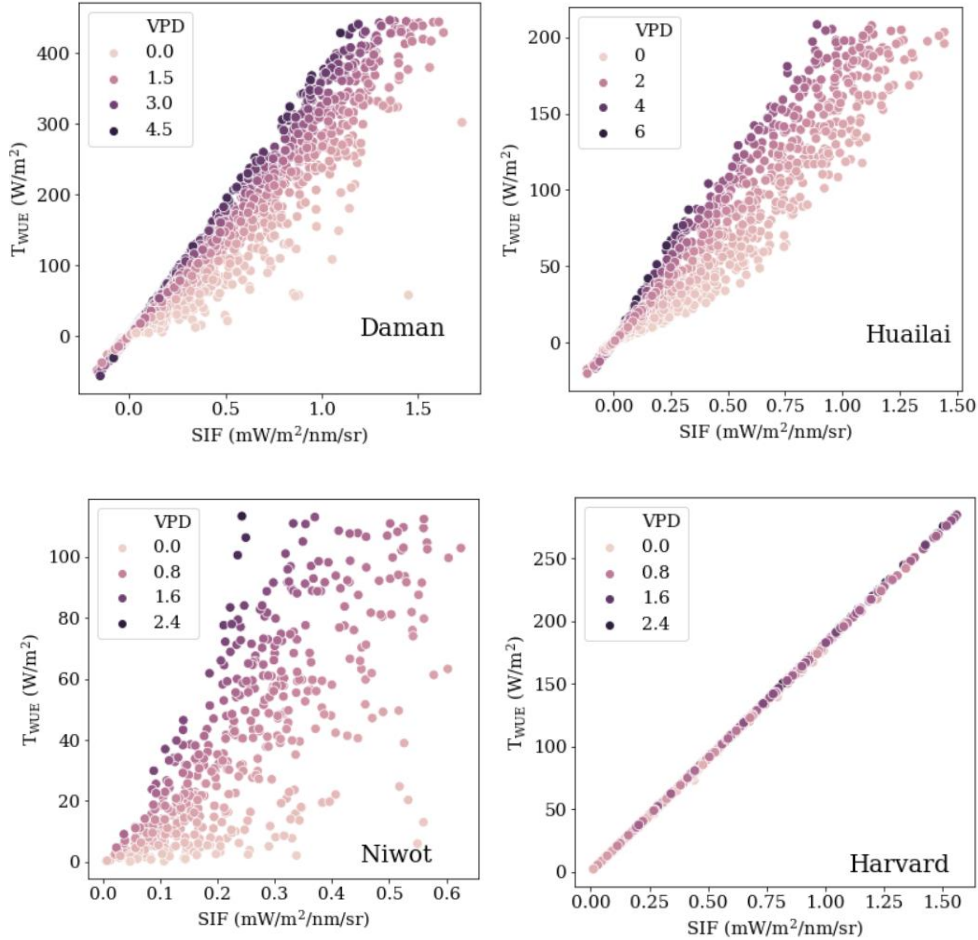
211 The reference T_{Zhou} may have considerable uncertainty. Plants do not always keep a specific
 212 response (square root) to VPD like described in Zhou's method. Moreover, in some ecosystems,
 213 the soil evaporation can not be ignored even in the peak growing season [Li et al., 2019; Stoy et

214 *al.*, 2019]. Here we also compared the LE derived by three SIF-T models with LE observed by
215 eddy covariance in Fig1 (lower row). The linear model, WUE model and conductance model have
216 $R^2 = 0.56, 0.61$ and 0.73 and $RMSE = 94.60, 91.20$ and 68.07 W/m^2 respectively. Same as
217 compared with T_{Zhou} , the conductance model outperforms the other two models with the highest
218 coefficients of determination and the lowest RMSE. The WUE model also has better performance
219 than the widely used linear model.

220 The performance of three models at different sites is shown in Table 4. When compared to
221 T_{Zhou} , the conductance model shows the best performance at DM and HF sites with much higher
222 $R^2 = 0.84, 0.62$, and lower $RMSE = 47.98, 77.43$ W/m^2 respectively. The linear model shows the
223 best performance with $R^2 = 0.58$ and $RMSE = 50.94$ W/m^2 at the HL site. The WUE model
224 outperforms other models at the NR site. When compared to LE observed by eddy covariance, the
225 conductance model shows outstanding performance at all four sites. Besides, the WUE model has
226 the lowest $RMSE = 67.38$ W/m^2 at the NR site.

227 3.2 Sensitivity analysis of variables in two models

228 To explore what influences the relationship between SIF and T, we analyzed the sensitivity
229 of variables in WUE and conductance models. For the WUE model, the scatterplot between SIF
230 and T_{WUE} is shown in Fig. T is the product of SIF and VPD in the WUE model, which means SIF
231 and VPD interact with each other closely and the effect of the SIF is modified by the VPD. The
232 parameter k_5 for four sites are 0.33, 0.45, 0.97, and 0.02 respectively. With the increase of VPD,
233 the slopes of SIF-T get steeper and points get denser, while in the low VPD condition, the points
234 are relatively sparse, which indicates the relationship between SIF and T is more linear under high
235 VPD. Especially we find the relation between SIF and T at the HF site is not sensitive to the VPD.



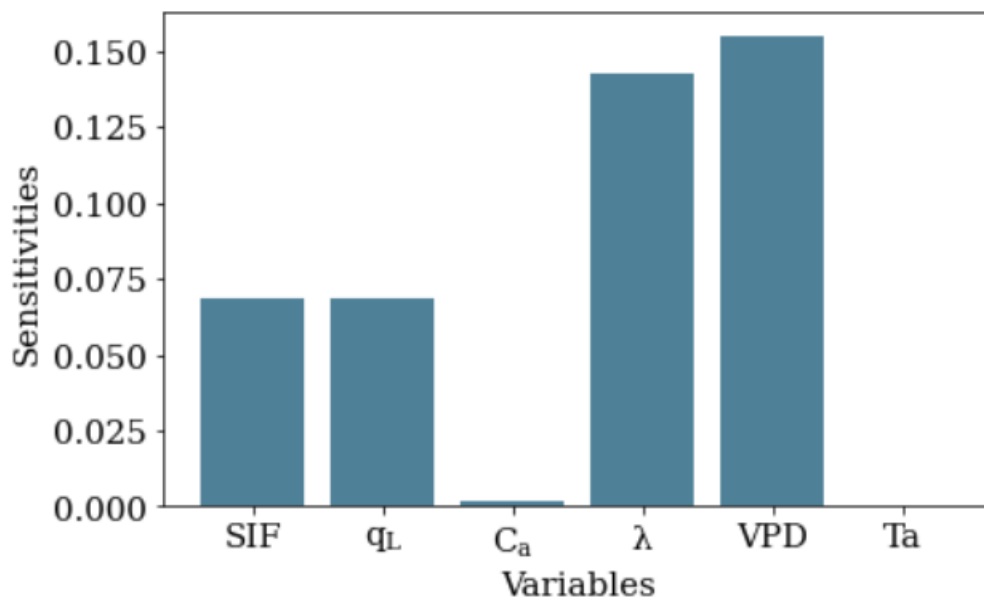
236

237

238 **Figure 2.** Scatter plot between SIF and T modeled by the WUE method. The vapor pressure deficit
 239 (VPD) has the unit kPa.

240 Atmospheric dryness is also important in modeling the stomatal behavior by SIF. For C4
 241 plants, relative humidity is used to describe the response of stomatal conductance to air dryness in
 242 the empirical Ball Berry model. For C3 plants, we investigated the sensitivities of different
 243 variables in the g_s model \mathcal{F} by RBD-FAST-Random Balance Designs Fourier Amplitude
 244 Sensitivity Test [Tarantola et al., 2006]. According to Fig 3, \mathcal{F} is sensitive to VPD, λ , SIF, and
 245 q_L . VPD exhibits the highest first-order sensitivities with values equaling 0.15, which is much
 246 higher than SIF with value 0.07. The marginal water use efficiency λ also plays an important role
 247 in \mathcal{F} with sensitivity value equaling 0.14. The fraction of open Photosynthesis II reaction centers
 248 q_L is as important as SIF (1st-order sensitivity: 0.07), which is due to the electron transport rate J

249 is the product of SIF and q_L (Eqn 1). What's more, the g_s model is not sensitive to air temperature
 250 and ambient carbon dioxide concentration. However, VPD used here is calculated by temperature
 251 and relative humidity of the air. The temperature information contained in VPD (R^2 of T_a -VPD is
 252 0.48) may impair the role of air temperature in the model. In this paper, q_L is calculated by a simple
 253 empirical equation of PAR. In fact, q_L is also related to the dark reaction but poorly studied [Baker,
 254 2008; Gu *et al.*, 2019]. More researches about q_L will improve our understanding of the
 255 relationship between SIF and J , further the SIF-GPP and SIF-T relationships.



256
 257 **Figure 3.** Sensitivity analysis of variables in the stomatal conductance model \mathcal{F} of C3 plants. The
 258 height of bars shows the 1st order sensitivities of different variables.

259 **4 Discussion**

260 Above all, the assumptions of carbon-water coupling may affect our results. Comparing with
 261 the linear model, though the influence of VPD on water-use efficiency is included in the WUE
 262 model, soil moisture, hydraulic conductance, and other environmental variables can influence
 263 water-use efficiency independently [Leakey *et al.*, 2019; Lin *et al.*, 2015; Liu *et al.*, 2020]. For the
 264 conductance model, plants under stress or competition tend to change the optimal stomatal
 265 conductance behavior [Wolf *et al.*, 2016]. The carbon-water economy is also influenced by traits

266 of plants and environmental variables [*Bloom et al.*, 1985; *Buckley et al.*, 2017]. This concept is
267 out of the scope of this paper, but a more physiologically based water-carbon coupling framework
268 will improve the models from the bottom up.

269 Due to the absence of direct measurements of transpiration during the research period, three
270 models were calibrated by the water budget balance framework, which might introduce uncertainty
271 to parameters, further the performance of the models. Net radiation is separated into energy
272 intercepted by canopy and soil available energy by 1D Beer's law. The simple structure of Beer's
273 law can introduce great uncertainty, especially for heterogeneity canopy and plants with leaves
274 highly anisotropic leaves in the azimuthal direction [*Ponce De León and Bailey*, 2019]. In the
275 conductance model, canopy available energy is also used to estimate transpiration in the two-
276 source Penman-Monteith model. So, the conductance model is more sensitive to the partition of
277 energy, in other words, suffers more uncertainty from Beer's law. Uncertainty of input data might
278 also affect our results. Firstly, VPD in the air was used to assess the aridity stress by assuming the
279 canopy and the atmosphere are fully coupled. Yet, ecosystems (DM, HL, and HF) with dense
280 closed canopy tend to decouple from the air [*De Kauwe et al.*, 2017; *Li et al.*, 2019; *Lin et al.*,
281 2018]. As both the WUE model and conductance model shows great sensitivity to VPD, using
282 VPD at the leaf scale could help to improve the performance of the SIF-T relationship. Secondly,
283 SIF data from four sites are measured by different instruments and derived by different methods
284 as mentioned above. Moreover, the FOV and height of the observation systems vary among sites.
285 So, it is difficult to derive universal parameters for all sites. Last but not the least, the canopy-scale
286 SIF was directly used to model transpiration due to data restriction. Nevertheless, recent papers
287 indicated the relationship between SIF and GPP is strongly affected by the structure of the canopy

288 [Liu *et al.*, 2019; Zeng *et al.*, 2019]. We suggest downscaling of SIF from the canopy scale to the
289 photosystem scale may improve the performance of the model.

290 Data from four sites were used to test the performance of the models. It is inadequate to show
291 the real potential of two models. With more and more in-situ observations from different
292 ecosystems, understanding of the underlying mechanism between SIF and T will be deepened.
293 Recently, SIF products with higher temporal-spatial resolution from different satellites, different
294 bands [Du *et al.*, 2018; Köhler *et al.*, 2018; Köhler *et al.*, 2020], and derivative products based on
295 machine learning [Li and Xiao, 2019; Ma *et al.*, 2020; Yu *et al.*, 2019; Zhang *et al.*, 2018] became
296 available. WUE model and conductance model can be easily combined with remote sensing ET/T
297 model like TSEB [Norman *et al.*, 1995; Song *et al.*, 2016] and PML ([Zhang *et al.*, 2019] or used
298 for assimilating SIF into a land surface model. With these data and models in this paper, estimating
299 T via SIF at the big scale becomes promising.

300 **5 Conclusion**

301 In this study, in-situ hourly SIF and corresponding meteorological variables, eddy covariance
302 observation and vegetation indexes at C3 and C4 sites were collected. Two SIF-T models based
303 on water-use efficiency and stomatal conductance were developed and tested upon this data. Both
304 models outperformed simple linear analysis with higher R^2 and lower RMSE. Our results indicate
305 the SIF-T relationship depends on air dryness. These two carbon-water coupling models can be
306 easily combined with state-of-the-art remote sensing models or land process models. Moreover,
307 with the emergency of high temporal-spatial resolution SIF data, SIF will not only be a powerful
308 proxy for carbon flux but also water flux at the planetary scale.

309

310

311 **Table 1.** Summary of the stations used to build models.

	Latitude	Longitude	Period	Land Cover	Reference
DM	100.37°E	38.85°N	2017.6 – 2017.9; 2018.6 – 2018.9	Maize (C4)	[Liu et al., 2018; Liu et al., 2019]
HL	115.78°E	40.33°N	2017.7 – 2017.10; 2018.7 – 2018.10	Maize (C4)	[Liu et al., 2013; Liu et al., 2019]
NR	105.55°W	40.03°N	2017.6-2017.9; 2018.6-2018.7	Mixed temperate forest (C3)	[Burns et al., 2015; Magney et al., 2019]
HF	72.17° W	42.54°N	2013.6-2013.11	Evergreen needle leaf forest(C3)	[Munger W, 2020; Yang et al., 2015]

312

313 **Table 2.** Input and intermediate variables. AWS denotes auto weather station. FAO indicates the
314 computation methods of the variable is from <http://www.fao.org>.

Variables	Unit	Description	Source	Remarks
C_a	$\mu\text{mol/mol}$	ambient CO2 concentration	Observation	Eddy covariance
C_p	J/Kg/K	specific heat of air	1013	FAO
LEs	W/m^2	Soil evaporation	$\frac{f\Delta(R_n - A_c)}{\Delta + \gamma}$	[Fisher et al., 2008]
f	-	Soil evaporation constraint	$SM/(SM_{\max} - SM_{\min})$	-
ga	m/s	Aerodynamic conductance	$\frac{1}{v/(u^*)^2 + 6.2(u^*)^{-2/3}}$	[Monteith and Unsworth, 2013]
GPP _{ob}	$\mu\text{mol/m}^2/\text{s}$	Gross primary production	Separated from NEE observed by eddy covariance	[Lasslop et al., 2010; Reichstein et al., 2005]
LE	W/m^2	Latent heat	Observation	Eddy covariance
PAR	W/m^2	photosynthetically active radiation	Observation	AWS
P	kPa	Air Pressure	Observation	AWS
q _L	-	Fraction of open Photosynthesis II reaction centers	$\exp(-\beta \text{PAR})$	This paper
Rh	-	Relative humidity	Observation	AWS
Rn	W/m^2	Net radiation	Observation	AWS
SIF	$\text{mW/m}^2/\text{sr/nm}$	Sun-induced chlorophyll fluorescence	Observation	-
SM	%	Soil moisture	Observation	Thermal Dissipation Probe
Ta	°C	Air temperature	Observation	AWS

u^*	m/s	Friction Velocity	Observation	Eddy covariance
v	m/s	Wind speed	Observation	AWS
VPD	kPa	Vapor pressure deficit	$(100 - Rh)/100 \times 0.6108 \times \exp(17.27 Ta/(Ta + 237.3))$	FAO
Δ	kPa/K	Slope of saturation vapor pressure curve	$(2503 \exp(17.27 Ta/(Ta + 237.3))) / ((Ta + 237.3)^2)$	FAO
γ	kPa/K	psychrometric constant	$0.665 \times 0.001P$	FAO
Γ	$\mu\text{mol/mol}$	CO2 compensation point in the absence of mitochondrial	$36.9 + 1.18(Ta - 25) + 0.036(Ta - 25)^2$ for C3; 0 for C4	[Katul et al., 2010]
ρ_a	kg/m^3	Air density	$1.292 - 0.00428 Ta$	FAO

315

316 **Table 3.** Parameters needed to be calibrated.

317

	Parameter	Lower	Upper	Reference
Linear model	K1	5	50	[Liu et al., 2017; Sun et al., 2017]
	K2	3	50	[Huang et al., 2016]
WUE model	K3	5	50	[Liu et al., 2017; Sun et al., 2017]
	K4	3	30	[Zhou et al., 2015]; This paper
	K5	0	1	[Lin et al., 2018]
Conductance model (C4)	β	0	0.001	This paper; [Gu et al., 2019]
	a	10	300	This paper
	m	2.5	8.8	[Miner et al., 2017]
Conductance model (C3)	β	0	0.01	This paper; [Gu et al., 2019; Katul et al., 2010]
	a	10	300	This paper
	λ	10	200	[Cowan and GD, 1977; Katul et al., 2010]

318

319 **Table 4.** Coefficient of determination (R^2) and root mean square error (RMSE) at different sites.
320 Best values are marked with the bold font.

	Reference	DM		HL		NR		HF	
		R^2	RMSE	R^2	RMSE	R^2	RMSE	R^2	RMSE
T_{linear}	T_{Zhou}	0.46	104.25	0.58	50.94	0.49	35.32	0.45	70.66
T_{WUE}		0.55	99.11	0.55	53.66	0.69	29.26	0.45	71.80

T_{gs}	0.84	47.98	0.57	54.79	0.46	41.27	0.62	77.43	
LE_{linear}	0.56	119.97	0.52	67.69	0.27	68.69	0.44	81.92	
LE_{WUE}	LE	0.55	114.20	0.55	66.71	0.36	67.38	0.44	84.87
LE_{gs}		0.87	69.36	0.60	64.53	0.40	68.82	0.55	70.71

321

322

References:

- 323 Alemohammad, S. H., et al. (2017), Water, Energy, and Carbon with Artificial Neural Networks (WECANN): a
 324 statistically based estimate of global surface turbulent fluxes and gross primary productivity using solar-induced
 325 fluorescence, *BIOGEOSCIENCES*, 14(18), 4101-4124.
- 326 Bai, Y., et al. (2019), Quantifying plant transpiration and canopy conductance using eddy flux data: An underlying
 327 water use efficiency method, *AGR FOREST METEOROL*, 271, 375-384.
- 328 Baker, N. R. (2008), Chlorophyll Fluorescence: A Probe of Photosynthesis In Vivo, *ANNU REV PLANT BIOL*, 59(1),
 329 89-113.
- 330 Ball, J. T., et al. (1987), A model predicting stomatal conductance and its contribution to the control of photosynthesis
 331 under different environmental conditions, in *Progress in photosynthesis research*, edited, pp. 221-224, Springer.
- 332 Beer, C., et al. (2009), Temporal and among-site variability of inherent water use efficiency at the ecosystem level,
 333 *GLOBAL BIOGEOCHEM CY*, 23(2), n/a-n/a.
- 334 Berry, J. A., et al. (2010), Stomata: key players in the earth system, past and present, *CURR OPIN PLANT BIOL*,
 335 13(3), 232-239.
- 336 Bloom, A. J., et al. (1985), Resource limitation in plants-an economic analogy, *Annual review of Ecology and*
 337 *Systematics*, 16(1), 363-392.
- 338 Buckley, T. N., et al. (2017), Optimal plant water economy, *Plant, Cell & Environment*, 40(6), 881-896.
- 339 Burns, S. P., et al. (2015), The influence of warm-season precipitation on the diel cycle of the surface energy balance
 340 and carbon dioxide at a Colorado subalpine forest site, *BIOGEOSCIENCES*, 12(23), 7349-7377.
- 341 Cowan, I. R., and F. GD (1977), Stomatal function in relation to leaf metabolism and environment.
- 342 Damm, A., et al. (2015), Far-red sun-induced chlorophyll fluorescence shows ecosystem-specific relationships to
 343 gross primary production: An assessment based on observational and modeling approaches, *REMOTE SENS*
 344 *ENVIRON*, 166, 91-105.
- 345 De Kauwe, M. G., et al. (2017), Ideas and perspectives: how coupled is the vegetation to the boundary layer?
 346 *BIOGEOSCIENCES*, 14(19), 4435-4453.
- 347 Du, S., et al. (2018), Retrieval of global terrestrial solar-induced chlorophyll fluorescence from TanSat satellite, *SCI*
 348 *BULL*, 63(22), 1502-1512.
- 349 Duan, Q., et al. (1994), Optimal use of the SCE-UA global optimization method for calibrating watershed models, *J*
 350 *HYDROL*, 158(3-4), 265-284.
- 351 Fisher, J. B., et al. (2008), Global estimates of the land - atmosphere water flux based on monthly AVHRR and
 352 ISLSCP-II data, validated at 16 FLUXNET sites, *REMOTE SENS ENVIRON*, 112(3), 901-919.
- 353 Frankenberg, C., and J. Berry (2018), Solar induced chlorophyll fluorescence: Origins, relation to photosynthesis and
 354 retrieval.
- 355 Gentine, P., et al. (2019), Coupling between the terrestrial carbon and water cycles—a review, *ENVIRON RES LETT*,
 356 14(8), 83003.
- 357 Good, S. P., et al. (2015), WATER RESOURCES. Hydrologic connectivity constrains partitioning of global terrestrial
 358 water fluxes, *SCIENCE*, 349(6244), 175-177.
- 359 Gorelick, N., et al. (2017), Google Earth Engine: Planetary-scale geospatial analysis for everyone, *REMOTE SENS*
 360 *ENVIRON*, 202, 18-27.
- 361 Gu, L., et al. (2019), Sun-induced Chl fluorescence and its importance for biophysical modeling of photosynthesis
 362 based on light reactions, *NEW PHYTOL*, 223(3), 1179-1191.
- 363 Guanter, L., et al. (2014), Global and time-resolved monitoring of crop photosynthesis with chlorophyll fluorescence,
 364 *Proceedings of the National Academy of Sciences*, 111(14), E1327-E1333.
- 365 Herman, J., and W. Usher (2017), SALib: an open-source Python library for sensitivity analysis, *Journal of Open*
 366 *Source Software*, 2(9), 97.
- 367 Houska, T., et al. (2015), SPOTting model parameters using a ready-made python package, *PLOS ONE*, 10(12),
 368 e145180.
- 369 Huang, M., et al. (2019), Air temperature optima of vegetation productivity across global biomes, *NAT ECOL EVOL*,
 370 3(5), 772-779.

Non-peer reviewed EarthArXiv preprint

【This manuscript has not been submitted to any journal and not under any peer review yet】

- 371 Jasechko, S., et al. (2013), Terrestrial water fluxes dominated by transpiration, *NATURE*, 496(7445), 347-350.
372 Jonard, F., et al. (2020), Value of sun-induced chlorophyll fluorescence for quantifying hydrological states and fluxes:
373 Current status and challenges, *AGR FOREST METEOROL*, 291, 108088.
374 Katul, G., et al. (2010), A stomatal optimization theory to describe the effects of atmospheric CO₂ on leaf
375 photosynthesis and transpiration, *ANN BOT-LONDON*, 105(3), 431-442.
376 Köhler, P., et al. (2018), Global Retrievals of Solar - Induced Chlorophyll Fluorescence With TROPOMI: First
377 Results and Intersensor Comparison to OCO - 2, *GEOPHYS RES LETT*, 45(19), 10, 410-456, 463.
378 Köhler, P., et al. (2020), Global Retrievals of Solar - Induced Chlorophyll Fluorescence at Red Wavelengths With
379 TROPOMI, *GEOPHYS RES LETT*, 47(15).
380 Kool, D., et al. (2014), A review of approaches for evapotranspiration partitioning, *AGR FOREST METEOROL*, 184,
381 56-70.
382 Lasslop, G., et al. (2010), Separation of net ecosystem exchange into assimilation and respiration using a light response
383 curve approach: critical issues and global evaluation, *GLOBAL CHANGE BIOL*, 16(1), 187-208.
384 Leakey, A., et al. (2019), Water Use Efficiency as a Constraint and Target for Improving the Resilience and
385 Productivity of C3 and C4 Crops, *ANNU REV PLANT BIOL*, 70, 781-808.
386 Leuning, R., et al. (2008), A simple surface conductance model to estimate regional evaporation using MODIS leaf
387 area index and the Penman-Monteith equation, *WATER RESOUR RES*, 44(10).
388 Li, X., et al. (2018), Solar-induced chlorophyll fluorescence is strongly correlated with terrestrial photosynthesis for
389 a wide variety of biomes: First global analysis based on OCO-2 and flux tower observations, *GLOBAL CHANGE*
390 *BIOL*, 24(9), 3990-4008.
391 Li, X., et al. (2019), A simple and objective method to partition evapotranspiration into transpiration and evaporation
392 at eddy-covariance sites, *AGR FOREST METEOROL*, 265, 171-182.
393 Li, X., and J. Xiao (2019), A Global, 0.05-Degree Product of Solar-Induced Chlorophyll Fluorescence Derived from
394 OCO-2, MODIS, and Reanalysis Data, *REMOTE SENS-BASEL*, 11(5), 517.
395 Lin, C., et al. (2018), Diel ecosystem conductance response to vapor pressure deficit is suboptimal and independent
396 of soil moisture, *AGR FOREST METEOROL*, 250-251, 24-34.
397 Lin, Y., et al. (2015), Optimal stomatal behaviour around the world, *NAT CLIM CHANGE*, 5(5), 459-464.
398 Liu, L., et al. (2017), Directly estimating diurnal changes in GPP for C3 and C4 crops using far-red sun-induced
399 chlorophyll fluorescence, *AGR FOREST METEOROL*, 232, 1-9.
400 Liu, S., et al. (2018), The Heihe Integrated Observatory Network: A Basin-Scale Land Surface Processes Observatory
401 in China, *VADOSE ZONE J*, 17(1), 180021-180072.
402 Liu, S. M., et al. (2013), Measurements of evapotranspiration from eddy-covariance systems and large aperture
403 scintillometers in the Hai River Basin, China, *J HYDROL*, 487, 24-38.
404 Liu, X., et al. (2019), Downscaling of solar-induced chlorophyll fluorescence from canopy level to photosystem level
405 using a random forest model, *REMOTE SENS ENVIRON*, 231, 110772.
406 Liu, X., et al. (2019), Atmospheric Correction for Tower-Based Solar-Induced Chlorophyll Fluorescence Observations
407 at O₂-A Band, *REMOTE SENS-BASEL*, 11(3), 355.
408 Liu, X., et al. (2020), Improving the potential of red SIF for estimating GPP by downscaling from the canopy level to
409 the photosystem level, *AGR FOREST METEOROL*, 281, 107846.
410 Liu, Y., et al. (2020), Plant hydraulics accentuates the effect of atmospheric moisture stress on transpiration, *NAT*
411 *CLIM CHANGE*, 10(7), 691-695.
412 Lu, X., et al. (2018), Potential of solar-induced chlorophyll fluorescence to estimate transpiration in a temperate forest,
413 *AGR FOREST METEOROL*, 252, 75-87.
414 Ma, Y., et al. (2020), Generation of a Global Spatially Continuous TanSat Solar-Induced Chlorophyll Fluorescence
415 Product by Considering the Impact of the Solar Radiation Intensity, *REMOTE SENS-BASEL*, 12(13), 2167.
416 Maes, W. H., et al. (2020), Sun-induced fluorescence closely linked to ecosystem transpiration as evidenced by
417 satellite data and radiative transfer models, *REMOTE SENS ENVIRON*, 249, 112030.
418 Magney, T. S., et al. (2019), Mechanistic evidence for tracking the seasonality of photosynthesis with solar-induced
419 fluorescence, *Proceedings of the National Academy of Sciences*, 201900278.
420 Miner, G. L., et al. (2017), Estimating the sensitivity of stomatal conductance to photosynthesis: a review, *Plant, Cell*
421 *& Environment*, 40(7), 1214-1238.
422 Monteith, J., and M. Unsworth (2013), Principles of environmental physics: plants, animals, and the atmosphere.
423 Munger W, W. S. (2020), Canopy-Atmosphere Exchange of Carbon, Water and Energy at Harvard Forest EMS Tower
424 since 1991. Harvard Forest Data Archive: HF004., edited.
425 Myneni, R., et al. (2015), MCD15A3H MODIS/Terra+ Aqua Leaf Area Index/FPAR 4-day L4 Global 500m SIN Grid

- 426 V006., edited.
- 427 Norman, J. M., et al. (1995), Source approach for estimating soil and vegetation energy fluxes in observations of
428 directional radiometric surface temperature, *AGR FOREST METEOROL*, 77(3-4), 263-293.
- 429 Pagán, B., et al. (2019), Exploring the Potential of Satellite Solar-Induced Fluorescence to Constrain Global
430 Transpiration Estimates, *REMOTE SENS-BASEL*, 11(4), 413.
- 431 Ponce De León, M. A., and B. N. Bailey (2019), Evaluating the use of Beer's law for estimating light interception in
432 canopy architectures with varying heterogeneity and anisotropy, *ECOL MODEL*, 406, 133-143.
- 433 Reichstein, M., et al. (2005), On the separation of net ecosystem exchange into assimilation and ecosystem respiration:
434 review and improved algorithm, *GLOBAL CHANGE BIOL*, 11(9), 1424-1439.
- 435 Ryu, Y., et al. (2019), What is global photosynthesis? History, uncertainties and opportunities, *REMOTE SENS*
436 *ENVIRON*, 223, 95-114.
- 437 Schimel, D., et al. (2019), Flux towers in the sky: global ecology from space, *NEW PHYTOL*, 224(2), 570-584.
- 438 Scott, R. L., and J. A. Biederman (2017), Partitioning evapotranspiration using long-term carbon dioxide and water
439 vapor fluxes, *GEOPHYS RES LETT*, 44(13), 6833-6840.
- 440 Shan, N., et al. (2019), Modeling canopy conductance and transpiration from solar-induced chlorophyll fluorescence,
441 *AGR FOREST METEOROL*, 268, 189-201.
- 442 Song, L., et al. (2016), Applications of a thermal-based two-source energy balance model using Priestley-Taylor
443 approach for surface temperature partitioning under advective conditions, *J HYDROL*, 540, 574-587.
- 444 Stoy, P. C., et al. (2019), Reviews and syntheses: Turning the challenges of partitioning ecosystem evaporation and
445 transpiration into opportunities, *BIOGEOSCIENCES*, 16(19), 3747-3775.
- 446 Sun, Y., et al. (2017), OCO-2 advances photosynthesis observation from space via solar-induced chlorophyll
447 fluorescence, *SCIENCE*, 358(6360), m5747.
- 448 Tarantola, S., et al. (2006), Random balance designs for the estimation of first order global sensitivity indices, *RELIAB*
449 *ENG SYST SAFE*, 91(6), 717-727.
- 450 Tol, C., et al. (2014), Models of fluorescence and photosynthesis for interpreting measurements of solar - induced
451 chlorophyll fluorescence, *Journal of Geophysical Research: Biogeosciences*, 119(12), 2312-2327.
- 452 Way, D. A., et al. (2014), Increasing water use efficiency along the C3 to C4 evolutionary pathway: a stomatal
453 optimization perspective, *J EXP BOT*, 65(13), 3683-3693.
- 454 Wolf, A., et al. (2016), Optimal stomatal behavior with competition for water and risk of hydraulic impairment,
455 *Proceedings of the National Academy of Sciences*, 113(46), E7222-E7230.
- 456 Wolz, K. J., et al. (2017), Diversity in stomatal function is integral to modelling plant carbon and water fluxes, *NAT*
457 *ECOL EVOL*, 1(9), 1292-1298.
- 458 Yang, X., et al. (2015), Solar-induced chlorophyll fluorescence that correlates with canopy photosynthesis on diurnal
459 and seasonal scales in a temperate deciduous forest, *GEOPHYS RES LETT*, 42(8), 2977-2987.
- 460 Yang, Y., et al. (2015), An analytical model for relating global terrestrial carbon assimilation with climate and surface
461 conditions using a rate limitation framework, *GEOPHYS RES LETT*, 42(22), 9825-9835.
- 462 Yu, L., et al. (2019), High - Resolution Global Contiguous SIF of OCO - 2, *GEOPHYS RES LETT*, 46(3), 1449-1458.
- 463 Yuan, W., et al. (2007), Deriving a light use efficiency model from eddy covariance flux data for predicting daily
464 gross primary production across biomes, *AGR FOREST METEOROL*, 143(3-4), 189-207.
- 465 Zeng, Y., et al. (2019), A practical approach for estimating the escape ratio of near-infrared solar-induced chlorophyll
466 fluorescence, *REMOTE SENS ENVIRON*, 232, 111209.
- 467 Zhang, Y., et al. (2014), Estimation of vegetation photosynthetic capacity from space-based measurements of
468 chlorophyll fluorescence for terrestrial biosphere models, *GLOBAL CHANGE BIOL*, 20(12), 3727-3742.
- 469 Zhang, Y., et al. (2018), A global spatially contiguous solar-induced fluorescence (CSIF) dataset using neural
470 networks, *BIOGEOSCIENCES*, 15(19), 5779-5800.
- 471 Zhang, Y., et al. (2019), Coupled estimation of 500 m and 8-day resolution global evapotranspiration and gross
472 primary production in 2002 - 2017, *REMOTE SENS ENVIRON*, 222, 165-182.
- 473 Zhang, Z., et al. (2020), Reduction of structural impacts and distinction of photosynthetic pathways in a global
474 estimation of GPP from space-borne solar-induced chlorophyll fluorescence, *REMOTE SENS ENVIRON*, 240, 111722.
- 475 Zhou, S., et al. (2014), The effect of vapor pressure deficit on water use efficiency at the subdaily time scale,
476 *GEOPHYS RES LETT*, 41(14), 5005-5013.
- 477 Zhou, S., et al. (2015), Daily underlying water use efficiency for AmeriFlux sites, *Journal of Geophysical Research:*
478 *Biogeosciences*, 120(5), 887-902.
- 479 Zhou, S., et al. (2018), Water use efficiency and evapotranspiration partitioning for three typical ecosystems in the
480 Heihe River Basin, northwestern China, *AGR FOREST METEOROL*, 253-254, 261-273.

Non-peer reviewed EarthArXiv preprint

【This manuscript has not been submitted to any journal and not under any peer review yet】

481