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# VESIcal Part I: An open-source thermodynamic model engine for mixed volatile (H<sub>2</sub>O-CO<sub>2</sub>) solubility in silicate melts

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## Key Points:

- The first comprehensive volatile solubility tool capable of processing large datasets automatically
- Seven built-in solubility models, with automatic calculation and plotting functionality
- Built in python and easily usable by scientists with any level of coding skill

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## Abstract

Thermodynamics has been fundamental to the interpretation of geologic data and modeling of geologic systems for decades. However, more recent advancements in computational capabilities and a marked increase in researchers' accessibility to computing tools has outpaced the functionality and extensibility of currently available modeling tools. Here we present VESIcal (Volatile Equilibria and Saturation Identification calculator): the first comprehensive modeling tool for H<sub>2</sub>O, CO<sub>2</sub>, and mixed (H<sub>2</sub>O-CO<sub>2</sub>) solubility in silicate melts that: a) allows users access to seven of the most popular models, plus easy inter-comparison between models; b) provides universal functionality for all models (e.g., functions for calculating saturation pressures, degassing paths, etc.); c) can process large datasets (1,000's of samples) automatically; d) can output computed data into an Excel spreadsheet or CSV file for simple post-modeling analysis; e) integrates plotting capabilities directly within the tool; and f) provides all of these within the framework of a python library, making the tool extensible by the user and allowing any of the model functions to be incorporated into any other code capable of calling python. The tool is presented within this manuscript, which is a Jupyter notebook containing worked examples accessible to python users with a range of skill levels. The basic functions of VESIcal can also be accessed via a web app (<https://vesical.anvil.app>). The VESIcal python library is open-source and available for download at <https://github.com/kaylai/VESIcal>.

## Plain Language Summary

Geologists use numerical models to understand and predict how volcanoes behave during storage (pre-eruption), eruption, and the composition and amount of volcanic gas released into the atmosphere of Earth and other planets. Most models are made by performing experiments on a limited dataset and creating a model that applies to that dataset. Some models combine lots of these individual models to make a generalized model that can apply to lots of different volcanoes. Many of these different models exist, and they all have specific uses, limitations, and pitfalls. Here we present the first tool, VESIcal, which acts as a simple interface to seven of the most commonly used models. VESIcal is written in python, so users can use VESIcal as an application or include it in their own models. VESIcal is the first tool that allows geologists to model thousands of data points automatically and provides a simple platform to compare results from different models in a way never before possible.

## 1 Introduction

Understanding the solubility and degassing of volatiles in silicate melts is a crucial component of modeling volcanic systems. As dissolved components, volatiles (primarily H<sub>2</sub>O and CO<sub>2</sub>) affect magma viscosity, rheology, and crystal growth. In addition, due to the strong dependence of volatile solubility on pressure, measured volatile concentrations in preserved high-pressure melts (i.e., melt inclusions: liquid magma trapped within crystals at high pressure, then brought to the surface during an eruption) can be used to determine pre-eruptive magmatic storage pressures, and thus depths. Importantly, volatile exsolution-driven overpressure of a magmatic system is likely the trigger of many explosive volcanic eruptions (Blake, 1984; Stock, Humphreys, Smith, Isaia, & Pyle, 2016; Tait, Jaupart, & Vergnolle, 1989). Once triggered, further drops in magmatic pressure caused by ascent of magma within a volcanic conduit result in the continuous exsolution of volatiles from the melt. Volatile elements experience a large positive volume change when moving from a dissolved to exsolved free fluid state. This expansion fuels a dramatic increase in the

66 magma’s buoyancy, which can often lead to a runaway effect in which the ascent and  
67 degassing of volatile-bearing magma eventually erupts at the surface in an explosive  
68 fashion. Working in concert with seismic and gas monitoring data, pre-eruptive mag-  
69 matic volatile concentrations as well as solubility and degassing modelling can be  
70 used in forensic and sometimes in predictive scenarios, helping us to understand and  
71 potentially mitigate the effects of explosive eruptions.

72 All of these processes depend directly on the solubility – or the capacity of a  
73 magma to hold in solution – of volatile elements. Over the last several decades, a  
74 veritable explosion of new volatile solubility data has opened the door to a plethora  
75 of models to describe the solubility of H<sub>2</sub>O, CO<sub>2</sub>, or mixed H<sub>2</sub>O-CO<sub>2</sub> fluid in mag-  
76 mas covering a wide compositional, pressure, and temperature range. Volatile solu-  
77 bility is highly dependent upon the composition of the host magma, making already  
78 challenging experiments more onerous to perform to encapsulate the range of mag-  
79 mas seen in nature. The most fundamental models (Dixon, Stolper, & Holloway,  
80 1995; Moore, Vennemann, & Carmichael, 1998; Stolper, 1982) focus on a specific  
81 range of magma bulk compositions (e.g., basalt or rhyolite only). Later studies filled  
82 in compositional gaps, some with an increased focus on mixed-volatile (H<sub>2</sub>O-  
83 CO<sub>2</sub>) studies, increasing the natural applicability of our models to more systems  
84 (Iacono-Marziano, Morizet, Trong, & Gaillard, 2012; Iacovino, Moore, Roggensack,  
85 Oppenheimer, & Kyle, 2013; Liu, Zhang, & Behrens, 2005). To date, there have  
86 been only a few significant efforts to create a holistic thermodynamic model cali-  
87 brated by a wide range of data in the literature. The most popular are MagmaSat  
88 (the mixed-volatile solubility model built into the software package MELTS v. 1.2.0;  
89 Ghiorso & Gualda, 2015) and the model of Papale, Morretti, and Barbato (2006).  
90 Both of these studies have made their source code available; the Papale et al. (2006)  
91 FORTRAN source code (titled Solwcad), web app, and a Linux program can be  
92 found at <http://www.pi.ingv.it/progetti/eurovolc/>, and very recently Mag-  
93 maSat has been made accessible via the ENKI thermodynamic python framework  
94 (<http://enki-portal.org/>).

95 Despite this communal wealth of solubility models, quantitative calculations of  
96 volatile solubility, and by extension saturation pressures, equilibrium fluid composi-  
97 tions, and degassing paths, remains a time-consuming endeavor. Modeling tools that  
98 are available are typically unable to process more than one sample at a time, requir-  
99 ing manual entry of the concentrations of 8-10 major oxides, temperature, as well as  
100 CO<sub>2</sub> and H<sub>2</sub>O concentrations to calculate saturation pressures, or  $X_{H_2O}$  to calculate  
101 dissolved volatile contents. This is particularly problematic for melt inclusion stud-  
102 ies, where saturation pressures are calculated for hundreds of inclusions, each with  
103 different entrapment temperatures, CO<sub>2</sub>, H<sub>2</sub>O, and major element concentrations.  
104 For example, the saturation pressures from 105 Gakkel ridge melt inclusions cal-  
105 culated in MagmaSat by Bennett, Jenner, Millet, Cashman, and Lissenberg (2019)  
106 required the manual entry of 1,365 values! The potential for user error in this data  
107 entry stage should not be overlooked.

108 In many cases, newly published solubility models do not include an accompa-  
109 nying tool, requiring users to correctly combine and interpret the relevant equations  
110 (e.g., Dixon, 1997; Dixon et al., 1995; Liu et al., 2005; Shishkina et al., 2014). This  
111 is problematic from a perspective of reproducibility of the multitude of studies  
112 utilizing these models, especially given that some of the equations in the original  
113 manuscripts contain typos or formatting errors. For some models, an excel spread-  
114 sheet was provided, or available at request from the authors. For example, Newman  
115 and Lowenstern (2002) included a simplified version of the Dixon (1997) model as  
116 part of “VolatileCalc”, which was written in Visual Basic for Excel. Due to its sim-  
117 plicity, allowing users to calculate saturation pressures, degassing paths, isobars and  
118 isopleths with a few button clicks and pop-up boxes, this tool has proved extremely

119 popular (with 766 citations at the time of writing). However, to calculate saturation  
 120 pressures using VolatileCalc, the user must individually enter the SiO<sub>2</sub>, H<sub>2</sub>O, CO<sub>2</sub>  
 121 content and temperature of every single sample into pop-up boxes. Similarly, the ex-  
 122 cel spreadsheet for the Moore et al. (1998) model calculates dissolved H<sub>2</sub>O contents  
 123 based on the concentration of 9 oxides, temperature, and the fraction of X<sub>H<sub>2</sub>O</sub> in the  
 124 vapor, which must be pasted in for every sample. Finally, Allison, Roggensack, and  
 125 Clarke (2019) provide an excel spreadsheet that allows users to calculate fugacities,  
 126 partial pressures, isobars, isopleths and saturation pressures. Again, parameters for  
 127 each sample must be entered individually, with no way to calculate large numbers of  
 128 samples automatically.

129 Some of these published models and tools are at risk of being lost to time,  
 130 since spreadsheet tools (particularly earlier studies published before journal-  
 131 provided hosting of data and electronic supplements was commonplace) must  
 132 be obtained by request to the author. Even if the files are readily available, pro-  
 133 grams used to open and operate them may not support deprecated file formats.  
 134 More recently, authors have provided web-hosted interfaces to calculating sat-  
 135 uration pressures and dissolved volatile contents (e.g., Iacono-Marziano et al.  
 136 2012; <http://calcul-isto.cnrs-orleans.fr/>, and Ghiorso and Gualda 2015;  
 137 [http://melts.ofm-research.org/CORBA\\_CTserver/GG-H2O-CO2.html](http://melts.ofm-research.org/CORBA_CTserver/GG-H2O-CO2.html)). Ghiorso  
 138 and Gualda (2015) also provide a Mac application. While more accessible in the  
 139 present time, this does not negate the issue of the longevity of these models. The  
 140 link provided in the Iacono-Marziano et al. (2012) manuscript returns an error “this  
 141 site cannot be reached”, although email contact with the author directed us towards  
 142 the newer link given above. Similarly, the link to the H<sub>2</sub>O-CO<sub>2</sub> equation of state  
 143 web calculator that Duan and Zhang (2006) provided in their manuscript returns a  
 144 404 error.

145 While we certainly advocate for the continued refinement of solubility models,  
 146 including the completion of new experiments in poorly studied yet critical compo-  
 147 sitional spaces such as andesites (Wieser, Iacovino, Moore, Matthews, & Allison,  
 148 submitted), a perhaps more crucial step at this juncture is in the development of a  
 149 tool that can apply modern computational solutions to making our current knowl-  
 150 edge base of volatile solubility in magmas accessible and enduring.

151 Here we present VESIcal (Volatile Equilibria and Saturation Identification  
 152 calculator): a python-based thermodynamic volatile solubility model engine that  
 153 incorporates seven popular volatile solubility models under one proverbial roof. The  
 154 models included in VESIcal are (also see Table 1):

- 155 1. MagmaSat: VESIcal’s default model. The mixed-volatile solubility model  
 156 within MELTS v. 1.2.0 (Ghiorso & Gualda, 2015)
- 157 2. Dixon: The simplification of the Dixon (1997) model as implemented in  
 158 VolatileCalc (Newman & Lowenstern, 2002)
  - 159 • DixonWater and DixonCarbon are available as pure-fluid models
- 160 3. MooreWater: (Moore et al. 1998; water only, but H<sub>2</sub>O fluid concentration can  
 161 be specified)
- 162 4. Liu: (Liu et al., 2005)
  - 163 • LiuWater and LiuCarbon are available as pure-fluid models
- 164 5. IaconoMarziano: (Iacono-Marziano et al., 2012)
  - 165 • IaconoMarzianoWater and IaconoMarzianoCarbon are available as pure-  
 166 fluid models
- 167 6. ShishkinaIdealMixing: (Shishkina et al., 2014) using pure-H<sub>2</sub>O and pure-CO<sub>2</sub>  
 168 models and assuming ideal mixing. In general, the pure-fluid versions of this  
 169 model should be used

- 170 • ShishkinaWater and ShishkinaCarbon are available as pure-fluid models
- 171 7. AllisonCarbon: (Allison et al. 2019, carbon only)
- 172 (a) AllisonCarbon-vesuvius (default; phonotephrite from Vesuvius, Italy)
- 173 (b) AllisonCarbon-sunset (alkali basalt from Sunset Crater, AZ, USA)
- 174 (c) AllisonCarbon-sfvf (basaltic andesite from San Francisco Volcanic Field,
- 175 AZ, USA)
- 176 (d) AllisonCarbon-erebus (phonotephrite from Erebus, Antarctica)
- 177 (e) AllisonCarbon-etna (trachybasalt from Etna, Italy)
- 178 (f) AllisonCarbon-stromboli (alkali basalt from Stromboli, Italy)

179 As any individual model is only valid within its calibrated range (see below), and  
 180 each model is parameterized and expressed differently (e.g., empirical vs. thermo-  
 181 dynamic models), it is impractical to simply combine them into one large model.  
 182 Instead, VESICAL is a single tool that can access and utilize all of these models, with  
 183 an extensive pressure-temperature-composition calibration range (Fig. 1). VESICAL  
 184 represents the first volatile solubility tool with the ability to perform calculations  
 185 for multiple samples at once, with built-in functionality for extracting data from an  
 186 Excel or CSV file. In addition, the code is written such that it is flexible (sample,  
 187 calculation type, and model type can be chosen discreetly) and extensible (VESICAL  
 188 code can be imported for use in python scripts, and the code is formatted such that  
 189 new volatile models can be added).

190 Importantly, VESICAL has been designed for practicality and ease of use.  
 191 It is designed to be used by anyone, from someone who is completely unfamiliar  
 192 with coding to an adept programmer. The non-coder user can interact with  
 193 VESICAL through a webapp (<https://vesical.anvil.app>) or directly within this  
 194 manuscript, which utilizes the user-friendly Jupyter Notebook format, allowing them  
 195 to upload a file with data, execute the various example calculations provided be-  
 196 low, and save the results to an Excel or CSV file to work with outside of VESICAL.  
 197 This notebook also incorporates built-in plotting options for easy visualization of  
 198 user data and calculated results. More experienced programmers may wish to use  
 199 the more advanced functionality provided by VESICAL, including the ability to hy-  
 200 bridize models (e.g., use one model for H<sub>2</sub>O and another for CO<sub>2</sub>) or write their own  
 201 routines and code calling VESICAL methods. VESICAL is an open source tool and as  
 202 such is far less prone to the preservation issues discussed above. Because the VESI-  
 203 cal code is hosted on GitHub, every change to the code is tracked publicly (Perkel,  
 204 2016). VESICAL's current release (version 0.9.10) is also archived on Zenodo, which  
 205 provides a static citable DOI (10.5281/zenodo.4652839) for the current version of the  
 206 code, along with a snapshot of the GitHub repository at the time of release.

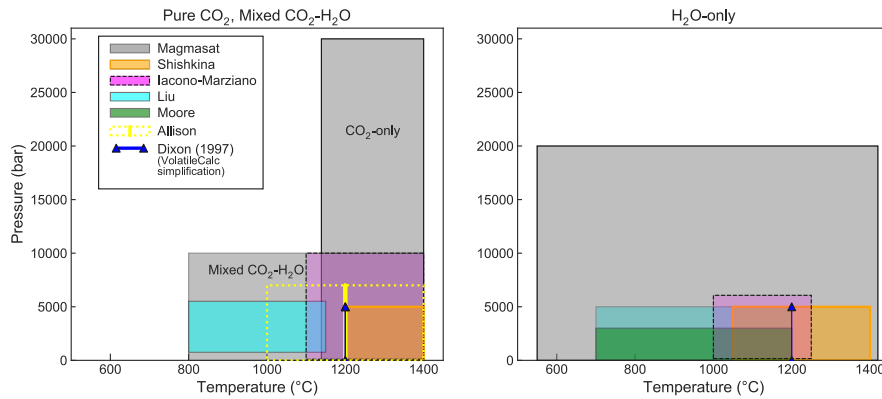
207 A detailed history of volatile solubility modeling and the implications of VESI-  
 208 cal are explored in detail in the companion manuscript to this work, Wieser et al.  
 209 (submitted).

## 210 **2 Research Methodology**

211 Navigating the array of models implemented in VESICAL can be challenging.  
 212 How can a user determine which model best suits their needs? MagmaSat (the de-  
 213 fault model in VESICAL) is the most widely calibrated in P-T-X space, and so we  
 214 recommend it for the majority of cases. Where a user wishes to use the other imple-  
 215 mented models, we provide some tools to help choose the most appropriate model  
 216 (see Supplement). These tools are described in more detail in Section 3.2 on compar-  
 217 ing user data to model calibrations.

**Table 1.** Calibration ranges of VESICAL models

Model/Reference	Species	P (bar)	T (°C)	Compositional range	Notes
<b>MagmaSat</b> <i>Ghiorso and Gualda, 2015</i>	H <sub>2</sub> O	0-20,000 <sup>1</sup>	550-1420 <sup>1</sup>	Very broad compositional range of natural silicate melts: subalkaline picobasalts to rhyolites, including a variety of mafic and silicic alkaline compositions	<sup>1</sup> Ranges extracted from Fig. 2d of Ghiorso and Gualda, 2015
	CO <sub>2</sub>	0-30,000 <sup>1</sup>	1139-1400 <sup>1</sup>		
<b>Dixon</b> <i>Simplification of Dixon (1997) used in VolatileCalc (Newman and Lowenstern, 2002)</i>	H <sub>2</sub> O - CO <sub>2</sub>	0-5000 <sup>2</sup>	800-1400 <sup>1</sup>	Alkali basalts: 40-49 wt% SiO <sub>2</sub>	<sup>1</sup> Warnings implemented in VolatileCalc (Newman and Lowenstern, 2002). <sup>2</sup> Calibration range suggested by Lesné et al. (2011) <sup>3</sup> Calibration range suggested by Iacono-Marziano et al. (2012) <sup>4</sup> Calibration temperature of Dixon (1997)
		0-2000 <sup>2</sup>	600-1500 <sup>1</sup>		
<b>MooreWater</b> <i>Moore et al. 1998</i>	H <sub>2</sub> O	0-3000 <sup>1</sup>	700-1200 <sup>1</sup>	Broad compositional range: subalkaline basalts to rhyolites, alkaline trachybasalts-andesites, foidites, phonolites	<sup>1</sup> Author-suggested calibration range. The calibration dataset spans 190 to 6067 bar, and 800-1200°C
<b>Liu</b> <i>Liu et al. 2005</i>	H <sub>2</sub> O - CO <sub>2</sub>	0-5000 <sup>1</sup>	700-1200 <sup>1</sup>	Haplogranites and rhyolites	<sup>1</sup> Author-suggested calibration range for the mixed fluid model. The calibration dataset covers 750-5510 bar and 800-1150°C for the Carbon model, and 1-5000 bar and 700-1200°C for the water model
<b>Iacono-Marziano</b> <i>Iacono-Marziano et al., 2012</i>	H <sub>2</sub> O - CO <sub>2</sub>	95-10,500 (mostly <5000) <sup>1</sup>	1100-1400 (preferably 1200-1300) <sup>2</sup>	Predominantly mafic compositions: subalkaline and alkaline basalts-andesites	<sup>1</sup> Range of calibration dataset, as authors do not specifically state a calibration range. We note that the vast majority of experiments were conducted at <5000 bar. <sup>2</sup> Authors state that most experiments were conducted between 1200-1300°C (whole range 1100-1400°C)
<b>Shishkina</b> <i>Shishkina et al. 2014</i>	H <sub>2</sub> O <sup>1</sup>	0-5000 <sup>2</sup>	1050-1400 (preferably 1150-1250) <sup>2,3</sup>	Mafic and intermediate compositions: Subalkaline basalts-basaltic andesites, alkali basanites-phonolites. SiO <sub>2</sub> <65 wt%.	<sup>1</sup> Although their empirical expressions are for pure fluids, they were mostly calibrated on mixed CO <sub>2</sub> -H <sub>2</sub> O experiments. <sup>2</sup> Author-suggested range <sup>3</sup> Note, this model contains no temperature term.
	CO <sub>2</sub> <sup>1</sup>	500-5000 <sup>2</sup>	1200-1250 <sup>2,3</sup>	Predominantly mafic compositions: subalkaline basalts, alkaline basanites, trachybasalts	
<b>AllisonCarbon</b> <i>Allison et al., 2019</i>	CO <sub>2</sub>	0-7000 <sup>1</sup>	1200 <sup>2</sup> (~1000-1400)	Alkali-rich mafic magmas from 6 volcanic fields. Separate model coefficients for each composition.	<sup>1</sup> Author-suggested range. The calibration dataset spans: (SPVF-4133-6141 bar, Sunset Crater; 4071-6098 bar, Erebus; 4078-6175 bar, Vesuvius; 269-6175 bar, Etna-455-6199, Stromboli-524-6080) <sup>2</sup> Note, all calculations performed at 1200 °C (the experimental temperature). Authors suggest results generally applicable between 1000-1400 °C



**Figure 1.** Illustrations showing the calibrated ranges of VESICAL models in pressure-temperature space. Due to difficulty in differentiating between pure-CO<sub>2</sub> and mixed fluid experiments in the literature, plots are subdivided into: experiments performed with pure-CO<sub>2</sub> or mixed (H<sub>2</sub>O-CO<sub>2</sub>) fluid; and pure-H<sub>2</sub>O fluid.

218 A list of model names recognized by VESICAL can be retrieved by executing  
 219 the command `v.get_model_names()`, assuming VESICAL has been imported as `v` as  
 220 is demonstrated in worked examples below. Note that the above model names are  
 221 given in terms of how to call them within VESICAL (e.g., `model='MooreWater'`).  
 222 Allison et al. (2019) provides unique model equations for each of the six alkali-rich  
 223 mafic magmas investigated in their study. The default model in VESICAL is that  
 224 calibrated for Vesuvius magmas, whose calibration has the widest pressure range of  
 225 the study (Table 1). Setting a model name of `'AllisonCarbon'` within VESICAL will  
 226 thus result in calculations using the `AllisonCarbon-vesuvius` model equations.

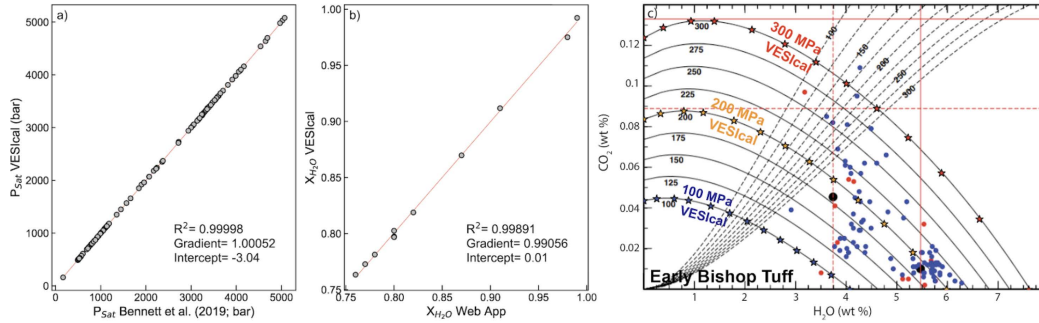
227 All of the calculations implemented in VESICAL can be performed using any of  
 228 the models included. The code is structured by calculation rather than by model,  
 229 which provides an intuitive way for users to interact with the code and compare  
 230 outputs from multiple models. A python method defined for each calculation takes  
 231 the model name and any applicable data as arguments and returns the results of  
 232 the calculation. Each method performs five key functions: 1) creates the requested  
 233 model object and performs any necessary pre-processing (e.g., ensuring relevant data  
 234 are present; normalizing data); 2) takes user input and performs the mathematical  
 235 calculation; 3) does any necessary processing of the output (e.g., normalizing totals);  
 236 4) checks that the model is being used within its calibrated range; and 5) returns  
 237 calculated outputs in an intuitive and manipulatable format (e.g., a python dictio-  
 238 nary, a figure, or a pandas DataFrame). Results of calculations can be saved to one  
 239 or more Excel or CSV files. To demonstrate that VESICAL returns results which are  
 240 comparable with pre-existing tools, we have performed a number of tests, which are  
 241 described in the Supplementary Information (Text S2).

## 242 2.1 Model Calibrations and Benchmarking

243 The pressure, temperature, and compositional calibration ranges of the seven  
 244 models implemented in VESICAL are shown in Table 1 and Figure 1. VESICAL abides  
 245 by statements of caution made by the authors of these models regarding their ex-  
 246 trapolation by informing the user if a calculation is being performed outside of a  
 247 model's calibrated range. In this case, the code returns a warning message, which  
 248 is as specific as possible, along with the requested output. We provide these cali-  
 249 brations along with several Jupyter notebooks in the supplementary material (Sup-  
 250 plementary Text S3-S4 and Jupyter Notebooks S1-S7), which allow users to plot  
 251 their data amongst the calibrations of the different models to assess their suitability  
 252 for less objective measures (also see Section 3.2). Detailed descriptions of the seven  
 253 solubility models implemented in VESICAL, including information about their calibra-  
 254 tion range in terms of melt composition, pressure, and temperature, are given in this  
 255 manuscript's companion paper Wieser et al. (submitted).

256 Testing was undertaken to ensure that VESICAL faithfully reproduces the re-  
 257 sults of all incorporated models. When possible, all models were benchmarked by  
 258 testing VESICAL outputs against those of a relevant published calculator (e.g., web  
 259 apps or Excel macros). The models of Shishkina et al. (2014) and Liu et al. (2005)  
 260 were published with no such tool and so testing instead compares VESICAL outputs  
 261 to experimental conditions or analyses and, where possible, plots VESICAL results  
 262 against published figures. All models underwent multiple tests, the results of which  
 263 are shown in the supplement (Supplementary Text S3-S4 and Supplemental Jupyter  
 264 Notebooks S1-S7). For all models, VESICAL reproduced the results from previous  
 265 tools (e.g., web apps, Excel spreadsheets) to within  $\pm 1\%$  relative and often on the  
 266 order of  $\pm 0.1\%$  relative.

267 MagmaSat, VESICAL's default model, underwent three tests, the results of  
 268 which are shown in Fig. 2: 1. Comparison of saturation pressures from MORB melt



**Figure 2.** Benchmarking of VESICAL against MagmaSat. a. Comparison of saturation pressures calculated with VESICAL against those by Bennett et al. (2019) using the MagmaSat app for Mac. Samples are all MORB melt inclusions, and pressures were calculated at a temperature unique to each sample. b. Equilibrium fluid compositions calculated with VESICAL against those calculated with the MagmaSat web app. c. Individual points along the 1,000, 2,000, and 3,000 bar isobars for the Early Bishop Tuff rhyolite calculated with VESICAL (stars) and plotted atop isobars published in Fig. 14 of Ghiorso and Gualda (2015).

269 inclusions in VESICAL to those published by Bennett et al. (2019), who used the  
 270 MagmaSat Mac App ( $R^2=0.99998$ ; Fig. 2a); 2. Comparison of fluid composition  
 271 ( $X_{H_2O}$ ) calculated with VESICAL and the web app ( $R^2=0.999$ , identical considering  
 272 the web app returns 2dp; Fig. 2b); 3. Comparison of isobars for the Early Bishop  
 273 Tuff calculated with VESICAL (star symbols) and isobars published in Fig. 14 of  
 274 Ghiorso and Gualda (2015) (Fig. 2c). VESICAL outputs using the model of Dixon  
 275 (1997) were tested against outputs from the VolatileCalc Excel spreadsheet (New-  
 276 man & Lowenstern, 2002) and a widely used Excel macro (see, e.g., Tucker et al.,  
 277 2019).

## 288 2.2 Format of the python library

279 In this section, the basic organization and use cases of VESICAL are discussed.  
 280 VESICAL relies heavily on python pandas, a python package designed for working  
 281 with tabulated data. Knowledge of pandas is not required to use VESICAL, and  
 282 we refer the user to the pandas documentation for an overview of the package  
 283 ([https://pandas.pydata.org/pandas-docs/stable/user\\_guide/index.html](https://pandas.pydata.org/pandas-docs/stable/user_guide/index.html)).

284 Specific details on how to perform model calculations are discussed in Section  
 285 3 and include worked examples. The VESICAL library is written so that users can  
 286 interact first and foremost with the calculation they want to perform. Five standard  
 287 calculations can be performed with any model in the library:

- 288 1. `calculate_dissolved_volatiles()`
- 289 2. `calculate_equilibrium_fluid_composition()`
- 290 3. `calculate_saturation_pressure()`
- 291 4. `calculate_isobars_and_isopleths()` (plus functionality for plotting; only  
 292 for mixed volatiles models)
- 293 5. `calculate_degassing_path()` (plus functionality for plotting; only for mixed  
 294 volatiles models).



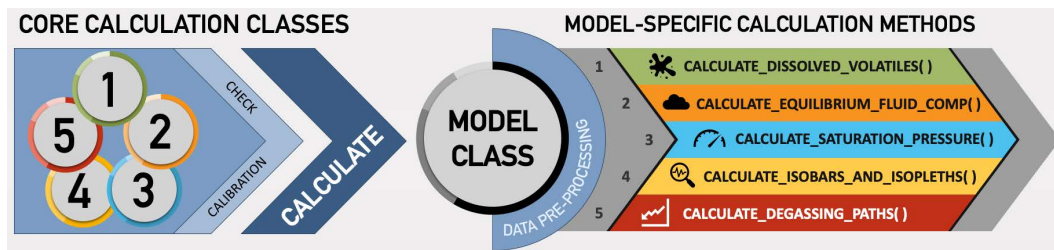
295 Fig. 3 illustrates the basic organization of the code. First, the user determines  
 296 which calculation they wish to perform by accessing one of the five core calculation  
 297 classes (listed above). In this step, the user specifies any input parameters needed  
 298 for the calculation (e.g., sample composition in wt% oxides, pressure in bars, tem-  
 299 perature in °C and fluid composition “X<sub>H<sub>2</sub>O<sup>fluid</sup>” in terms of XH<sub>2</sub>O<sup>fluid</sup>) as well as the  
 300 model they wish to use. The default model is MagmaSat, but the user may spec-  
 301 ify any model in the library. As an example, the code to calculate the saturation  
 302 pressure of some sample using the MagmaSat model would be written as:</sub>

```
303 calculate_saturation_pressure(sample=mysample, temperature=850.0).  
304 result
```

305 where `mysample` is a variable (python dictionary or pandas Series) containing the  
 306 composition of the sample in oxide wt%, and the temperature is given in °C. Exam-  
 307 ples on how to create such a variable are given in Section 3. If a different model is  
 308 desired, for example Dixon (1997), it can be passed as:

```
309 calculate_saturation_pressure(sample=mysample, temperature=850.0,  
310 model='Dixon').result
```

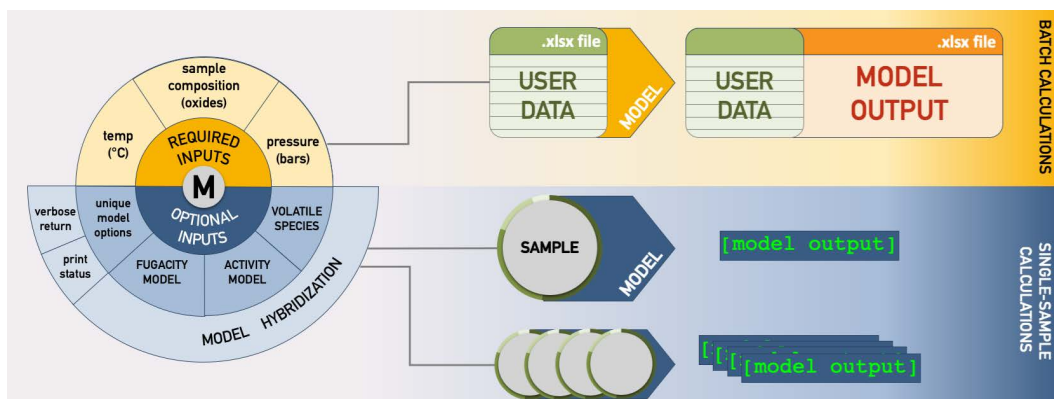
311 The core calculation classes each perform two functions: 1) a check is per-  
 312 formed to ensure that the user input is within the model’s recommended calibration  
 313 range; 2) the `calculate()` method sends the user input to the appropriate model.



**Figure 3.** Flowchart illustrating the basic organization of the python library. First, a user chooses a calculation to perform and calls one of the five core calculation classes. Here, any necessary parameters are passed such as sample composition, pressure, and temperature. A check is run to ensure the calculation is being performed within model-specified limits. The `Calculate()` class then calls on one of the `Model()` classes. The default model is MagmaSat, but a user may specify a different model when defining the calculation parameters. Standard pre-processing is then performed on the input data, and this pre-processing step is unique to each model. The processed data are then fed into a model-specific method to perform the desired core calculation.

314 Users can process individual samples (single-sample calculations) or entire  
 315 datasets (batch calculations; Fig. 4). If processing more than one sample, the “sim-  
 316 plest” way to interact with VESical is via batch calculations. Here, the user pro-  
 317 vides input data in the form of a Microsoft Excel spreadsheet (.xlsx file) or CSV file  
 318 and instructs the model to perform whatever calculation is desired. The model is  
 319 run on all samples and returns data formatted like a spreadsheet (using the python  
 320 pandas package), which contains the user’s original input data plus whatever model  
 321 outputs were calculated. The user can continue to work with returned data by sav-  
 322 ing the result to a variable (as is shown in all examples in this manuscript). Data  
 323 can then be exported to an Excel or CSV file with a simple command (see Section  
 324 3.10).

325 The syntax for processing a single sample is very similar to that for batch  
 326 calculations but provides the user direct access to more advanced features that can-  
 327 not be accessed via batch calculations (e.g., specifying fugacity or activity model,  
 328 hybridizing models; see Section 3.9). This also gives the user more flexibility in  
 329 integrating any VESICAL model function into some other python code.



**Figure 4.** Flowchart illustrating the different operational paths. On top, batch calculation is shown, in which an Excel or CSV file with any amount of samples is fed into the model, calculations are performed, and the original user data plus newly calculated values are returned and can be saved as an Excel or CSV file. Below, single-sample calculation is shown. These methods can run calculations on one sample at a time, but multi-sample calculations can be performed iteratively with code written by the user. Calculated values are returned as a variable. For single-sample calculations, more advanced modeling options can be set, and hybridization of models can be performed.

### 330 2.3 Running the code

331 VESICAL can be used in a number of ways: via this Jupyter notebook, via the  
 332 VESICAL web app, or by directly importing VESICAL into any python script.

333 VESICAL was born from functionality provided by ENKI and so all the files nec-  
 334 essary to use VESICAL are hosted on the ENKI server (<http://enki-portal.org/>).  
 335 A unique personal coding environment can be initiated by logging into the ENKI  
 336 production server using a GitLab username and password (which is free to obtain;  
 337 see directions on the ENKI website for specifics). The simplest way to use VESICAL  
 338 while retaining all of its functionality is within this very manuscript, in the form of  
 339 a Jupyter notebook. Because this manuscript and VESICAL python library files are  
 340 hosted on the ENKI server, code can be manipulated and executed in the code cells  
 341 below. Making changes won't affect the public version of this manuscript. Likewise,  
 342 any user can write their own python code using VESICAL by creating a Jupyter note-  
 343 book on the ENKI server and importing VESICAL as is demonstrated in the code  
 344 below.

345 Computation time on the ENKI server is limited by the server itself. VESI-  
 346 cal may run faster if installed locally. Advanced instructions on installing VESI-  
 347 cal on your own computer are provided in the Supplement (Supplementary Text  
 348 S1). Note that VESICAL requires installation of the ENKI thermoengine library to  
 349 function properly. Thermoengine is written in python but is based on the origi-  
 350 nal MELTS code (Ghiorso & Gualda, 2015; Ghiorso & Sack, 1995), which con-

351 tains MacOS-specific header files. The result is that thermoengine is most easily  
 352 installed on MacOS but can be installed on Windows and Linux operating sys-  
 353 tems via Docker (see thermoengine documentation for installation instructions;  
 354 <https://gitlab.com/ENKI-portal/ThermoEngine>).

355 The most limited but simplest method to interacting with VESICAL is through  
 356 the web app (<https://vesical.anvil.app>). The web app can currently perform  
 357 three of the five core calculations in batch process mode (via upload of an Excel or  
 358 CSV file). Some, but not all, optional parameters can be set.

359 **To run the code in this notebook**, nothing needs to be installed. Simply  
 360 execute the code cells below, changing parameters as desired. Custom data may be  
 361 processed by uploading an Excel or CSV file into the same folder containing this  
 362 notebook and then changing the filename in Section 3.1.

## 363 2.4 Documentation

364 This manuscript serves as an introduction to the VESICAL library aimed at  
 365 python users of all levels. However, the code itself is documented with explanations  
 366 of each method, its input parameters, and its returned values. This documentation  
 367 can be accessed at our readthedocs website (<https://vesical.readthedocs.io/>).  
 368 The documentation for any function can be viewed in a jupyter notebook by typ-  
 369 ing the function followed by a question mark and executing the cell (e.g., “`v.`  
 370 `calculate_saturation_pressure?`”).

371 Video tutorials are also available on the VESICAL YouTube ([https://](https://www.youtube.com/channel/UCpvCCs5KMXz0xXWm0seF8Qw)  
 372 [www.youtube.com/channel/UCpvCCs5KMXz0xXWm0seF8Qw](https://www.youtube.com/channel/UCpvCCs5KMXz0xXWm0seF8Qw)). Currently, the first  
 373 tutorial covers the basics of VESICAL. More videos for specific features and uses are  
 374 planned.

## 375 2.5 Generic methods for calculating mixed-fluid properties

376 VESICAL provides a set of methods for calculating the properties of mixed  
 377 CO<sub>2</sub>-H<sub>2</sub>O fluids, which can be used with any combination of H<sub>2</sub>O and CO<sub>2</sub> solu-  
 378 bility model. The use of generic methods allows additional models to be added to  
 379 VESICAL by defining only the (simpler) expressions describing pure fluid solubility.  
 380 Non-ideality of mixing in the fluid or magma phases can be incorporated by specify-  
 381 ing activity and fugacity models. A complete description of these methods, including  
 382 all relevant equations, can be found in the Supplement (Supplementary Text S2).

## 383 3 Workable example uses

384 In this section we detail how to use the various functions available in VESICAL  
 385 through worked examples. The python code presented below may be copied and  
 386 pasted into a script or can be edited and executed directly within the Jupyter note-  
 387 book version of this manuscript. For all examples, code in sections 3.0.2 and 3.1  
 388 must be executed to initialize the model and import data from the provided com-  
 389 panion Excel file. The following sections then may be executed on their own and do  
 390 not need to be executed in order.

391 In each example below, a generic “method structure” is given along with def-  
 392 initions of unique, required, and optional user inputs. The method structure is  
 393 simply for illustrative purposes and gives default values for every argument (input).  
 394 In some cases, executing the method structure as shown will not produce a sensible  
 395 result. For example, the default values for the `plot()` function (Section 3.8) contain

396 no data, and so no plot would be produced. Users should replace the default values  
397 shown with values corresponding to the samples or conditions of interest.

398 All examples will use the following sample data by default (but this can be  
399 changed by the user):

- 400 • Dataset from `example_data.xlsx` loaded in Section 3.1.1 (variable name `myfile`  
401 )
- 402 • Single composition defined in Section 3.1.2 (variable name `mysample`)
- 403 • Sample 10\* extracted from `example_data.xlsx` dataset in Section 3.1.3 (vari-  
404 able name `sample_10`)

405 Calculations performed on single samples or on a dataset imported from an  
406 Excel or CSV file containing many samples are executed in two distinct ways. Note  
407 that single sample calculations require that the argument `sample` be defined. To  
408 return the numerical result of the calculation, `.result` must be added to the end  
409 of the method, as shown below. Batch calculations are performed on the dataset  
410 itself, after that dataset is imported into VESICAL. Thus, the `sample` argument does  
411 not need to be defined discretely, since sample compositional information is stored  
412 within the dataset object. The two basic formats for performing calculations are:

413 *Single sample calculations*

```
414 myvariable = v.name_of_the_core_calculation(sample=mysample ,
415                                             argument1=value1, argument2
416                                             =value2).result
417
418
```

419 *Batch calculations*

```
420 myvariable = myfile.name_of_the_core_calculation(argument1=value1 ,
421                                                  argument2=value2)
422
423
```

424 where VESICAL has been imported as `v`, `myvariable` is some arbitrary  
425 variable name to which the user wishes to save the calculated output,  
426 `name_of_the_core_calculation` is one of the five core calculations, `mysample` is  
427 a variable containing compositional information in wt% oxides, `myfile` is a vari-  
428 able containing an `BatchFile` object created by importing an Excel or CSV file,  
429 and `argument1`, `argument2`, `value1`, and `argument2` are two required or optional  
430 arguments and their user-assigned values, respectively.

431 Workable examples detailed here are:

- 432 1. Loading, viewing, and preparing user data
  - 433 1.1 Loading a Batch file
  - 434 1.2 Defining a single sample composition
  - 435 1.3 Plotting user data
  - 436 1.4 Extracting a single sample from a Batch file
  - 437 1.5 Normalizing and transforming data
- 438 2. Calculating dissolved volatile concentrations
- 439 3. Calculating equilibrium fluid compositions
- 440 4. Calculating saturation pressures
- 441 5. Calculating and plotting isobars and isopleths
- 442 6. Calculating and plotting degassing paths

- 443 7. Plotting multiple calculations  
 444 8. Comparing results from multiple models  
 445 9. Model hybridization (Advanced)  
 446 10. Exporting data

### 447 **3.0.1 Function arguments and their definitions**

448 Each section below details what arguments are required or optional inputs and  
 449 gives examples of how to perform the calculations. Table 2 lists all arguments, both  
 450 required and optional, used in the five core calculations. Many of the function argu-  
 451 ments have identical form and use across all calculations, and so we list these here.  
 452 Any special cases are noted in the section describing that calculation.

453 The most commonly used arguments are:

454 **sample** *Single sample calculations only* The composition of a sample. A  
 455 VESICAL Sample object is created to hold compositional information about  
 456 sample. A Sample object can be created from a dictionary or pandas Series  
 457 containing values, with compositions of oxides in wt%, oxides in mol fraction,  
 458 or cations in mol fraction. This argument is not needed for batch calculations  
 459 since they are performed on BatchFile objects, which already contain sample  
 460 information. See examples for details.

461 **temperature, pressure, and X\_fluid:** the temperature in °C, the pressure in  
 462 bars, and the mole fraction of H<sub>2</sub>O in the H<sub>2</sub>O-CO<sub>2</sub> fluid, XH<sub>2</sub>O<sup>fluid</sup>. In all  
 463 cases, X\_fluid is optional, with a default value of 1 (pure H<sub>2</sub>O fluid). Note  
 464 that the X\_fluid argument is only used for calculation of dissolved volatile  
 465 concentrations.

466 *For single sample calculations*

467 Temperature, pressure, and X\_fluid should be specified as a numerical value.

468 *For batch calculations*

469 Temperature, pressure, and X\_fluid can either be specified as a numerical  
 470 value or as strings referring to the names of columns within the file containing  
 471 temperature, pressure, or X\_fluid values for each sample. If a numerical value  
 472 is passed for either temperature, pressure, or X\_fluid, that will be the value  
 473 used for one or all samples. If, alternatively, the user wishes to use tempera-  
 474 ture, pressure, and/or X\_fluid information in their BatchFile object, the title  
 475 of the column containing temperature, pressure, or X\_fluid data should be  
 476 passed in quotes (as a string) to **temperature**, **pressure**, and/or **X\_fluid**,  
 477 respectively. Note for batch calculations that if temperature, pressure, or  
 478 XH<sub>2</sub>O<sup>fluid</sup> information exists in the BatchFile but a single numerical value is  
 479 defined for one or both of these variables, both the original information plus  
 480 the values used for the calculations will be returned.

481 **verbose:** *Only for single sample calculations.* Always an optional argument  
 482 with a default value of False. If set to True, additional values of interest,  
 483 which were calculated during the main calculation, are returned in addition  
 484 to the results of the calculation.

485 `print_status`: *Only for batch calculations.* Always an optional argument,  
 486 which sometimes defaults to True and other times defaults to False (see  
 487 specific calculation section for details). If set to True, the progress of the cal-  
 488 culation will be printed to the terminal. The user may desire to see the status  
 489 of the calculation, as some calculations using MagmaSat can be somewhat  
 490 slow, particularly for large datasets.

491 `model`: Always an optional argument referring to the name of the desired  
 492 solubility model to use. The default is always “MagmaSat”.

**Table 2.** Matrix of all arguments used in the five core calculations, the nature of the argument (required or optional) and the input type or default value.

	dissolved_volatiles		equilibrium_fluid_comp		saturation_pressure		isobars_isopleths		degassing_path	
	SS	Batch	SS	Batch	SS	Batch	SS	Batch	SS	Batch
sample	wt% oxides		wt% oxides		wt% oxides		wt% oxides		wt% oxides	
temperature	°C	°C	°C	°C	°C	°C	°C	°C	°C	°C
pressure	bars	bars	bars	None						'saturation'
pressure_list										
X_fluid	1	1								
isopleth_list										
verbose	False		False		False		None			
model	'MagmaSat'	'MagmaSat'	'MagmaSat'	'MagmaSat'	'MagmaSat'	'MagmaSat'	'MagmaSat'	'MagmaSat'	'MagmaSat'	'MagmaSat'
print_status		True		False		True		True		
smooth_isobars								True		
smooth_isopleths								True		
fractionate_vapor								True		
init_vapor										0.0
										0.0

SS = Single-sample. Batch = batch processing. Color of cells corresponds to the type of argument: green=required; orange=optional; gray=argument not used. Values in cells indicate the unit or type of data to input for required arguments or the default value in the case of optional arguments.

### 3.0.2 Initialize packages

493 For any code using the VESICAL library, the library must be imported for use.  
 494 Here we import VESICAL as `v`. Any time we wish to use a function from VESICAL,  
 495 that function must be preceded by '`v.`' (e.g., `v.calculate_saturation_pressure`  
 496 ). Specific examples of this usage follow. Here we also import some other python  
 497 libraries that we will be using in the worked examples below.  
 498

#### Input

```

500 import VESICAL as v
501 import pandas as pd
502
503
504 #The following are options for formatting this manuscript
505 pd.set_option('display.max_colwidth', 0)
506 from IPython.display import display, HTML
507 %matplotlib inline
508
    
```

### 3.1 Loading, viewing, and preparing user data

509 All of the following examples will use data loaded in the code cells in this  
 510 section. Both batch processing of data loaded from a file and single-sample pro-  
 511 cessing are shown. An example file called 'example\_data.xlsx' is included with this  
 512 manuscript. You can load in your own data by first ensuring that your file is in the  
 513 same folder as this notebook and then by replacing the filename in the code cell  
 514 below with the name of your file. The code cell below must be executed for the  
 515 examples in the rest of this section to function properly.  
 516

### 3.1.1 Batch processing

Batch calculations are always facilitated via the `BatchFile()` class, which the user uses to specify the filename corresponding to sample data. Loading in data is as simple as calling `BatchFile(filename)`. Optionally, `units` can be used to specify whether the data are in wt% oxides, mol fraction oxides, or mol fraction cations. Calculations will always be performed and returned with melt composition in the default units (wt% oxides unless changed by the user) and fluid composition in mol fraction.

**Structure of the input file:** A file containing compositions (and optional pressure, temperature, or  $\text{XH}_2\text{O}^{\text{fluid}}$  information) on one or multiple samples can be loaded into VESICAL. The loaded file must be a Microsoft Excel file with the extension `.xls` or `.xlsx` or CSV file with the extension `.csv`. The file must be laid out in the same manner as the example file ‘example.data.xlsx’. The basic structure is also shown in Table 3.

Any extraneous columns that are not labeled as oxides or input parameters will be ignored during calculations. The first column titled ‘Label’ contains sample names. Note that the default assumption on the part of VESICAL is that this column will be titled ‘Label’. If no ‘Label’ column is found, the first non-oxide column name will be set as the index column, meaning this is how samples can be accessed by name (see Section 3.1.3). An index column can be specified by the user using the argument `label` (see documentation below). The following columns must contain compositional information as oxides. The only allowable oxides are:  $\text{SiO}_2$ ,  $\text{TiO}_2$ ,  $\text{Al}_2\text{O}_3$ ,  $\text{Fe}_2\text{O}_3$ ,  $\text{FeO}$ ,  $\text{Cr}_2\text{O}_3$ ,  $\text{MnO}$ ,  $\text{MgO}$ ,  $\text{CaO}$ ,  $\text{NiO}$ ,  $\text{CoO}$ ,  $\text{Na}_2\text{O}$ ,  $\text{K}_2\text{O}$ ,  $\text{P}_2\text{O}_5$ ,  $\text{H}_2\text{O}$ , and  $\text{CO}_2$ . Currently, VESICAL can only read these oxide names exactly as written (e.g., with no leading or trailing spaces and with correct capitalization), but functionality to interpret variations in how these oxides are entered is planned (e.g., such that “sio2. ” would be understood as “SiO2”). All of these oxides need not be included; if for example your samples contain no NiO concentration information, you can omit the NiO column. Omitted oxide data will be set to 0 wt% concentration. If other oxide columns not listed here are included in your file, they will be ignored during calculations. Notably, the order of the columns does not matter, as they are indexed by name rather than by position. Compositions can be entered either in wt% (the default), mol%, or mole fraction. If mol% or mole fraction data are loaded, this must be specified when importing the file.

Because VESICAL may misread column headings, we highly recommend that users examine their data after loading into VESICAL and before performing calculations. The user data, as it will be used by VESICAL, can be viewed at any time with `myfile.get_data()` (see generation of Table 3 below).

Pressure, temperature, or  $\text{XH}_2\text{O}^{\text{fluid}}$  data may optionally be included, if they are known. Column names for these data do not matter, as they can be specified by the user as will be shown in following examples.

The standard units used by VESICAL are always pressure in bars, temperature in  $^{\circ}\text{C}$ , melt composition as oxides in wt%, and fluid composition as mol fraction (typically specified as `X_fluid`, the mol fraction of  $\text{H}_2\text{O}$  in an  $\text{H}_2\text{O}$ - $\text{CO}_2$  fluid, ranging from 0-1). Sample compositions may be translated between wt%, mol fraction, and mol cations if necessary.

**Class structure:** `BatchFile(filename, sheet_name=0, file_type='excel', units='wtpt_oxides', label='Label', default_normalization='none', default_units='wtpt_oxides', dataframe=None)`

**Required inputs:**

567 **filename:** A file name must be passed in quotes. This file must be in the  
 568 same folder as the notebook or script that is calling it. This imports the data  
 569 from the file name given and saves it to a variable of your choosing.

570 **Optional inputs:** By default, the BatchFile class assumes that loaded data is in  
 571 units of wt%; alternatively, data in mol% or mole fraction may be loaded. In that  
 572 case, loaded data is converted into wt% values, since compositions must be in wt%  
 573 when performing model calculations.

574 **sheet\_name:** If importing data from an Excel file, this argument is used to  
 575 specify which sheet to import. Only one sheet can be imported to a single  
 576 BatchFile object. The default is '0', which imports the first sheet in the file,  
 577 regardless of its name.

578 **file\_type:** Specifies whether the file being imported is an Excel or CSV file.  
 579 This argument is never strictly necessary, as `BatchFile()` will automatically  
 580 detect whether an imported file is Excel or CSV if the file extension is one of  
 581 `.xls` or `.xlsx` (Excel) or `.csv` (CSV).

582 **units:** The units in which data are input. The default value is 'wtpt\_oxides'  
 583 for data as wt% oxides. The user can pass 'mol\_oxides' for data in mol frac-  
 584 tion oxides or 'mol\_cations' for data in mol fraction cations.

585 **default\_normalization:** The type of normalization to apply to the data  
 586 by default. One of: One of: `None`, `'standard'`, `'fixedvolatiles'`, or '  
 587 `additionalvolatiles'`. These normalization types are described in the  
 588 section on normalization below.

589 **default\_units:** The type of composition to return by default, one of: '  
 590 `wtpt_oxides'` (wt% oxides, default), `'mol_oxides'` (mol fraction oxides), or  
 591 `'mol_cations'` (mol fraction cations).

592 **label:** This is optional but can be specified if the column title referring to  
 593 sample names is anything other than "Label". The default value is "Label".  
 594 If no "Label" column is present and the label argument is not specified, the  
 595 first column whose first row is not one of VESICAL's recognized oxides will be  
 596 set as the index column. The index column will be used to select samples by  
 597 name.

598 **dataframe:** This argument is used for transforming a pandas DataFrame  
 599 object into a VESICAL BatchFile object. For convenience, this functionality is  
 600 also defined as a separate function `BatchFile_from_DataFrame(dataframe,`  
 601 `units='wtpt_oxides', label='Label')`.

## 602 **Outputs:**

603 A special type of python object defined in the VESICAL code known as an  
 604 BatchFile object.

605 **Input**



606  
607  
608

```
myfile = v.BatchFile('Supplement/Example_Datasets/example_data.xlsx')
```

609  
610  
611  
612  
613  
614  
615

Once the BatchFile object is created and assigned to a variable, the user can then access the data loaded from their file as `variable.get_data()`. In this example, the variable corresponding to the BatchFile object is named `myfile` and so the data in that file can be accessed with `myfile.get_data()`. Below, `myfile.get_data()` is saved to a variable we name `data`. The variable `data` is a pandas DataFrame object, which makes displaying the data itself quite simple and aesthetically pleasing, since pandas DataFrames mimic spreadsheets.

616  
617

Usage of `get_data()` allows the user to retrieve the data as originally entered or in any units and with any normalization supported by VESICAL.

618  
619

**Class structure:** `get_data(self, normalization=None, units=None, asBatchFile=False)`

620

**Optional inputs:**

621  
622

`normalization` or `units` may be passed, with options as defined in the description of BatchFile above.

623

`asBatchFile` Default is False. If True, will return a VESICAL BatchFile object.

624

**Outputs:**

625

A pandas dataframe or BatchFile object with all user data.

626

**Input**

627  
628  
629  
630

```
data = myfile.get_data()
data
```

**Table 3.** User input data: Compositions, pressures, and temperatures for several silicate melts as supplied in the file ‘example\_data.xlsx’

Label	CITATION	ROCK TYPE	SiO2	TiO2	Al2O3	Fe2O3	Cr2O3	FeO	MnO	MgO	NiO	CoO	CaO	Na2O	K2O	P2O5	H2O	CO2	Press	Temp
KIR3-6_1a	Tucker et al. (2019)	Basalt	48.248207	2.222114	11.692194	0.00	0.0	0.000000	0.079999	14.183817	0.0	0.0	9.892732	1.810522	0.352014	0.210479	0.424695	0.002873	62.5	1299.094712
KIR3-6_3a	Tucker et al. (2019)	Basalt	48.295691	2.165357	11.755584	0.00	0.0	0.000000	0.084045	13.403980	0.0	0.0	10.052578	2.268198	0.373328	0.204452	0.425984	0.006786	128.0	1283.419991
KIR3-6_4a	Tucker et al. (2019)	Basalt	49.124079	2.369984	12.172833	0.00	0.0	0.000000	0.098809	11.997899	0.0	0.0	10.308188	2.001863	0.396512	0.238996	0.437758	0.004984	124.5	1255.153759
10*	Roggensack (2001)	Basalt	47.960000	0.780000	18.770000	0.00	0.0	10.920000	0.150000	6.860000	0.0	0.0	12.230000	1.950000	0.210000	0.170000	4.500000	0.047900	4000.0	1200.000000
19*	Roggensack (2001)	Basalt	49.640000	0.710000	18.050000	0.00	0.0	10.540000	0.190000	6.430000	0.0	0.0	12.090000	1.990000	0.200000	0.170000	5.100000	0.111300	4000.0	1200.000000
25	Roggensack (2001)	Basalt	50.320000	0.720000	18.030000	0.00	0.0	10.110000	0.140000	5.650000	0.0	0.0	12.780000	1.800000	0.230000	0.230000	5.200000	0.043700	4000.0	1200.000000
SAT-M12-1	Moore et al. (1998)	Andesite	62.600000	0.630000	17.300000	2.01	0.0	2.010000	0.060000	2.650000	0.0	0.0	5.840000	4.050000	1.610000	0.240000	2.620000	0.000000	703.0	1100.000000
SAT-M12-2	Moore et al. (1998)	Andesite	62.600000	0.630000	17.300000	2.01	0.0	2.010000	0.060000	2.650000	0.0	0.0	5.840000	4.050000	1.610000	0.240000	5.030000	0.000000	1865.0	1100.000000
SAT-M12-4	Moore et al. (1998)	Andesite	62.600000	0.630000	17.300000	2.01	0.0	2.010000	0.060000	2.650000	0.0	0.0	5.840000	4.050000	1.610000	0.240000	6.760000	0.000000	2985.0	1050.000000
samp. P1968a	Myers et al. (2019)	Rhyolite	76.974880	0.855616	3.110636	0.00	0.0	4.788883	0.000000	12.549439	0.0	0.0	1.207910	0.138963	1.133094	0.000000	4.340000	0.007000	2000.0	900.000000
samp. P1968b	Myers et al. (2019)	Rhyolite	76.943845	0.133125	3.169657	0.00	0.0	4.763435	0.000000	12.446403	0.0	0.0	1.231728	0.140993	1.170806	0.000000	5.850000	0.012300	2000.0	900.000000
samp. P1968c	Myers et al. (2019)	Rhyolite	77.167205	0.119506	3.167827	0.00	0.0	4.814076	0.000000	12.229534	0.0	0.0	1.184773	0.138201	1.158924	0.000000	5.754571	0.010663	2000.0	900.000000
samp. HPR3-1_XL-3	Mercer et al. (2015)	Rhyolite	75.413866	0.095164	14.076992	0.00	0.0	0.654992	0.125882	0.012003	0.0	0.0	0.636124	3.703110	5.128392	0.000000	5.943750	0.010000	2000.0	0.000000
samp. HPR3-1_XL-4_INCL-1	Mercer et al. (2015)	Rhyolite	76.613586	0.095843	13.476762	0.00	0.0	0.620769	0.113495	0.032069	0.0	0.0	0.624350	3.676972	4.579799	0.000000	5.340000	0.008000	0.0	900.000000
AW-6	Iacovino et al. (2016)	Phonotephrite	48.030000	2.840000	18.120000	0.00	0.0	9.600000	0.230000	3.080000	0.0	0.0	7.570000	6.040000	3.080000	1.410000	1.420000	0.128900	2000.0	1050.000000
AW-46	Iacovino et al. (2016)	Basaltic-Trachyandesite	52.980000	2.180000	20.490000	0.00	0.0	5.540000	0.200000	2.000000	0.0	0.0	7.100000	5.680000	3.160000	0.660000	4.760000	0.343900	4000.0	1000.000000
KI-07	Iacovino et al. (2016)	Basanite	44.610000	4.370000	14.410000	0.00	0.0	10.600000	0.170000	7.690000	0.0	0.0	11.550000	3.930000	1.740000	0.820000	2.900000	0.113100	2000.0	1100.000000

631  
632  
633  
634  
635  
636

For the rest of this manuscript, data will be pulled from the example\_data.xlsx file (Supplemental Dataset S1), which contains compositional information for basalts (Roggensack, 2001; Tucker et al., 2019), andesites (Moore et al., 1998), rhyolites (Mercer et al., 2015; Myers, Wallace, & Wilson, 2019), and alkaline melts (phonotephrite, basaltic-trachyandesite, and basanite from Iacovino, Oppenheimer, Scaillet, and Kyle 2016). Several additional example datasets from the literature are

available in the Supplement (Supplementary Datasets S2-S5; Table 4). These include experimentally produced alkaline magmas from Iacovino et al. (2016, alkaline.xlsx), basaltic melt inclusions from Kilauea (Tucker et al., 2019) and Gakkel Ridge (Bennett et al., 2019, basalts.xlsx), basaltic melt inclusions from Cerro Negro volcano, Nicaragua (Roggensack, 2001, cerro\_negro.xlsx), and rhyolite melt inclusions from the Taupo Volcanic Center, New Zealand (Myers et al., 2019) and a topaz rhyolite from the Rio Grande Rift (Mercer et al., 2015, rhyolites.xlsx). Where available, the calibration datasets for VESICAL models are also provided (Supplementary Datasets S6-S7).

Input

```
pd.read_excel("Table_Example_Data.xlsx", index_col="Filename")
```

Output

**Table 4.** Example datasets included with VESICAL

Filename	Explanation	Compositions	Citations
example_data.xlsx	Example data used in this manuscript	Wide comp. range	Iacovino et al. (2016); Mercer et al. (2015); Myers et al. (2019); Roggensack (2001); Tucker et al. (2019)
alkaline.xlsx	Experimental glasses	Basanite to Tephriphonolite	Iacovino et al. (2016)
basalts.xlsx	Melt inclusion glasses	Basaltic	Tucker et al. (2019); Bennett et al. (2019)
cerro_negro.xlsx	Melt inclusion glasses	Basaltic	Roggensack (2001)
rhyolites.xlsx	Melt inclusion glasses	Rhyolitic	Mercer et al. (2015); Myers et al. (2019)

### 3.1.2 Defining a single sample

More advanced functionality of VESICAL is facilitated directly through the five core calculation classes. Each calculation requires its own unique inputs, but all calculations require that a sample composition be passed. We can pass in a sample either as a python dictionary or pandas Series. Below, we define a sample and name it `mysample`. Oxides are given in wt%. Only the oxides shown here can be used, but not all oxides are required. Any extra oxides (or other information not in the oxide list) the user defines will be ignored during calculations.

Much like is done to create a BatchFile object, we can create a VESICAL Sample object to represent our sample composition.

```
Class structure: Sample(composition, units='wtpt\_oxides',
default_normalization='none', default_units='wtpt\_oxides')
```

#### Required inputs:

**composition:** The composition of the sample in the format specified by the `units` parameter. The default is oxides in wt%.

#### Optional inputs:

`units`, `default_normalization`, and `default_units` have the same meaning here as in the BatchFile class described above.

#### Outputs:

A special type of python object defined in the VESICAL code known as a Sample object.

To manually input a bulk composition, fill in the oxides in wt% below:

Input

```
mysample = v.Sample({'SiO2': 77.3,
                    'TiO2': 0.08,
                    'Al2O3': 12.6,
                    'Fe2O3': 0.207,
                    'Cr2O3': 0.0,
                    'FeO': 0.473,
                    'MnO': 0.0,
                    'MgO': 0.03,
                    'NiO': 0.0,
                    'CoO': 0.0,
                    'CaO': 0.43,
                    'Na2O': 3.98,
                    'K2O': 4.88,
                    'P2O5': 0.0,
                    'H2O': 6.5,
                    'CO2': 0.05})
```

To see the composition of `mysample`, use the `get_composition(species=None, normalization=None, units=None, exclude_volatiles=False, asSampleClass=False)` method. By default, the composition is returned exactly as input above. `species` can be set as an element or oxide (e.g., "Si" or "SiO<sub>2</sub>") to return the float value for only that species. The composition can automatically be normalized using any of the standard normalization functions listed above and can be returned in any of the units discussed above. As with the `BatchFile.get_data()` function, a sample composition can be returned as a dictionary (default) or as a VESICAL Sample object (if `asSampleClass` is set to True).

The oxides considered by VESICAL are:

Input

```
print(v.oxides)
```

Output

```
['SiO2', 'TiO2', 'Al2O3', 'Fe2O3', 'Cr2O3', 'FeO', 'MnO', 'MgO', 'NiO', 'CoO', 'CaO', 'Na2O', 'K2O', 'P2O5', 'H2O', 'CO2']
```

### 3.1.3 Extracting a single sample from a batch file

Defined within the `BatchFile()` class, the method `get_sample_composition()` allows for the extraction of a melt composition from a loaded Excel or CSV file.

**Method structure:** `myfile.get_sample_composition(samplename, species=None, normalization=None, units=None, asSampleClass=False)`

**Required inputs:**

`samplename:` The name of the sample, as a string, as defined in the 'Label' column of the input file.

**Optional inputs:**

720 **species:** This is used if only the concentration of a single species (either  
721 oxide or element) is desired.

722 **normalization:** This is optional and determines the style of normaliza-  
723 tion performed on a sample. The default value is `None`, which returns the  
724 value-for-value un-normalized composition. Other normalization options are  
725 described in the `BatchFile` class description above.

726 **units:** The default is wt% oxides. Other options are described in the `Batch-`  
727 `File` class description above.

728 **asSampleClass:** Can be `True` or `False` (default). If set to `False`, this will  
729 return a dictionary with compositional values. If set to `True`, this will return  
730 a `Sample` object with compositional data stored within.

### 731 **Outputs:**

732 The bulk composition stored in a dictionary or `Sample` object.

### 733 **Input**

```
734 """To get composition from a specific sample in the input data:"""
735 sample_10 = myfile.get_sample_composition('10*', asSampleClass=True)
736
737
738 """To see the extracted sample composition, uncomment the line below by
739         removing the # and execute this code
740         cell"""
741 #sample_10.get_composition()
742
```

### 743 **3.1.4 Normalizing and transforming data**

744 Before performing model calculations on your data, it may be desired to nor-  
745 malize the input composition to a total of 100 wt%. For a user to decide whether  
746 normalization is prudent, is important to understand the influence any normal-  
747 ization, or lack thereof, to a composition will have on modeling results. Electron  
748 microprobe analyses of major elements in silicate glasses combined with volatile el-  
749 ement analyses by SIMS and FTIR often sum to less than 100 wt%. This deficiency  
750 is normally attributed to subsurface charging, matrix corrections, and unknown  
751 redox states of Fe and S during analyses by electron microprobe (see Huges et al.,  
752 2019). As an example, when normalized, a volatile-free basalt with a measured  $\text{SiO}_2$   
753 content of 46 wt% and an analytical total of 97 wt% actually contains 47.4 wt%  
754  $\text{SiO}_2$  (46/0.97; a 3% relative change in silica content). Many studies report major  
755 element data normalized to 100% with volatiles listed separately. The result is that,  
756 value for value, literature datasets can have totals several wt% less than 100 (if raw  
757 data are reported) or several wt% higher than 100 (if major elements are normalized  
758 anhydrous).

759 To deal with this variation, `VESICAL` provides users with four options for nor-  
760 malization. Normalization types are:

- 761 • `None` (no normalization)
- 762 • `'standard'`: Normalizes an input composition to 100%.

- ‘fixedvolatiles’: Normalizes major element oxides to 100 wt%, including volatiles. The volatile wt% will remain fixed, whilst the other major element oxides are reduced proportionally so that the total is 100 wt%.
- ‘additionalvolatiles’: Normalizes major element oxide wt% to 100%, assuming it is volatile-free. If H<sub>2</sub>O or CO<sub>2</sub> are passed to the function, their un-normalized values will be retained in addition to the normalized non-volatile oxides, summing to >100%.

Normalization can be performed on a Sample object or on all samples within a BatchFile object using the `get_composition()` or `get_data()` methods (e.g., `myfile.get_composition(normalization='standard')` or `mysample.get_composition(normalization='additionalvolatiles')`). Note that, since a BatchFile object may have other data in addition to sample compositions (e.g., information on pressure, temperature, other user notes), `BatchFile.get_composition()` returns only compositional data, whereas `BatchFile.get_data()` returns all data stored in the BatchFile object. The `normalization` argument can be passed to either. In the example below, we obtain the standard normalization of `mysample` and `myfile` and save these to new Sample and BatchFile objects called `mysample_normalized` and `myfile_normalized`. Note that `asSampleClass` or `asBatchFile` must be set to True in order to return a Sample or BatchFile object. Without this argument, a dictionary or pandas DataFrame will be returned and new Sample or BatchFile objects will need to be constructed from those in order to perform calculations on the normalized datasets.

#### Input

```

"""Retrieve the standard normalization for one sample"""
mysample_normalized = mysample.get_composition(normalization="standard",
                                              asSampleClass=True)
#print(mysample_normalized.get_composition())

"""Retrieve the standard normalization for all samples in a BatchFile"""
myfile_normalized = myfile.get_data(normalization="standard", asBatchFile=
                                   True)
#print(myfile_normalized.get_data())

```

The Liu and all six AllisonCarbon models are not sensitive to normalization because they contain no compositional terms. Similarly, the expressions for Shishkina and MooreWater contain compositional terms expressed solely in terms of anhydrous cation fractions; the `additionalvolatiles` and `fixedvolatiles` normalization routines do not affect the relative abundances of major elements (and therefore anhydrous cation fractions). Thus, Shishkina and MooreWater are only affected by the standard normalization routine. In contrast, the Dixon model is highly sensitive to the choice of normalization because its compositional term for both H<sub>2</sub>O and CO<sub>2</sub> is expressed solely in terms of the absolute melt SiO<sub>2</sub> content.

The expressions of Iacono-Marziano are parameterized in terms of hydrous cation fractions and NBO/O, and so this model is sensitive to `additionalvolatiles` or `fixedvolatiles` normalization routines, which will change the relative proportions of volatiles to major elements. Even so, the effect of normalization on volatile solubility calculations is relatively small and of similar magnitude to the discrepancy between the hydrous total and 100 for the hydrous model. Thus, the choice of normalization is only important when data has hydrous totals that differ significantly from 100%. The Iacono-Marziano web app normalizes input data a la VESICAL’s `additionalvolatiles` normalization routine. For consistency with the web app, VESICAL

815 automatically uses the additionalvolatiles normalization during calculations with this  
816 model.

817 The implementation of MagmaSat in VESICAL is sensitive to the relative pro-  
818 portion of major and volatile element components rather than the absolute concen-  
819 trations entered (as with the whole MELTS family of models). Thus, calculations  
820 using raw, fixed- and additionalvolatile routines yield different results. If the hy-  
821 drous total of an input composition is less than 100%, the fixedvolatile routine ef-  
822 fectively reduces the relative proportion of volatiles to major elements, so calculated  
823 saturation pressures go down. Conversely, if inputs have high hydrous totals, the  
824 fixedvolatile routine increases the relative proportion of volatiles in the system, so  
825 the saturation pressure goes up. As with Iacono-Marziano, the percent discrepancy  
826 between calculations for different normalization routines is similar to the difference  
827 between the total and 100%. For saturation pressure calculations, the MagmaSat  
828 app automatically normalizes input data a la VESICAL's fixedvolatiles routine. Thus,  
829 we suggest that users should normalize their inputs using fixedvolatiles for consis-  
830 tency with previous studies. However, for maximum flexibility, no normalization  
831 on inputs is forced and so must be set by the user if desired. Further discussion on  
832 the effect of normalization in MagmaSat is provided in Supporting Text S5 (and  
833 Supporting Figs S22-S26).

834 For example, consider a basalt with a measured SiO<sub>2</sub> content of 47.4 wt%,  
835 1000 ppm dissolved CO<sub>2</sub>, and an anhydrous (volatile-free) total of 96.77 wt%:

836 Input

```
837 mybasalt = v.Sample({'SiO2': 47.4,  
838                    'TiO2': 1.01,  
839                    'Al2O3': 17.46,  
840                    'Fe2O3': 0.89,  
841                    'FeO': 7.18,  
842                    'MgO': 7.63,  
843                    'CaO': 12.44,  
844                    'Na2O': 2.65,  
845                    'K2O': 0.03,  
846                    'P2O5': 0.08,  
847                    'CO2': 0.1})  
848  
849
```

850 We can apply each normalization routine to this sample and examine how this  
851 will affect the saturation pressure predicted by each model:

852 Input

```
853 """Normalize three ways"""  
854 mybasalt_std = mybasalt.get_composition(normalization="standard",  
855                                       asSampleClass=True)  
856 mybasalt_add = mybasalt.get_composition(normalization="additionalvolatiles",  
857                                       asSampleClass=True)  
858 mybasalt_fix = mybasalt.get_composition(normalization="fixedvolatiles",  
859                                       asSampleClass=True)  
860  
861  
862 """Choose a model to test"""  
863 mymodel = "IaconoMarziano"  
864  
865 for basalt, normtype in zip([mybasalt, mybasalt_std, mybasalt_add,  
866                             mybasalt_fix],
```

```

867         ["Raw", "standard", "additionalvolatiles", "
868                                         fixedvolatiles
869                                         "]):
870     print(str(normtype) +
871           " Saturation Pressure = " +
872           str(v.calculate_saturation_pressure(sample=basalt, temperature=1200
873                                               , model=mymodel).result))
874

```

## Output

```

875 Raw Saturation Pressure = 1848.031831425599
876 standard Saturation Pressure = 1906.5453789627868
877 additionalvolatiles Saturation Pressure = 1848.2673972122493
878 fixedvolatiles Saturation Pressure = 1848.2611364359402
879
880
881

```

882 Because the compositional effect on H<sub>2</sub>O solubility is smaller, so are the  
883 changes in calculated saturation pressures for a pure-H<sub>2</sub>O system, but they can  
884 still be significant for H<sub>2</sub>O-rich liquids (where high H<sub>2</sub>O contents can change totals  
885 and therefor SiO<sub>2</sub> contents more dramatically).

## 886 3.2 Comparing User Data to Model Calibrations: Which Model 887 Should I Use?

888 MagmaSat is the most thermodynamically robust model implemented in VESI-  
889 cal, and thus it is the most generally appropriate model to use (n.b. that it is also  
890 the most computationally expensive). However, one of the strengths of VESIcal is  
891 its ability to utilize up to seven different solubility models. Each of these models is  
892 based on its own calibration dataset, meaning the pressure-temperature-composition  
893 space over which models are calibrated is quite variable from model to model. The  
894 individual model calibrations are discussed in detail in this manuscript's companion  
895 paper (VESIcal Part II; Wieser et al., in prep).

896 For the remainder of this section, all example calculations are carried out with  
897 MagmaSat, the default model of VESIcal. To use any other VESIcal model, simply  
898 add 'model=' and the name of the desired model in quotes to any calculation (e.g.,  
899 `v.calculate_dissolved_volatiles(temperature=900, pressure=1000, model`  
900 `= "Dixon")`). The model names recognized by VESIcal are: MagmaSat, Shishki-  
901 naIdealMixing, Dixon, IaconoMarziano, Liu, AllisonCarbon, and MooreWater. For  
902 more advanced use cases such as hybridizing models (see Section 3.9), pure-H<sub>2</sub>O and  
903 pure-CO<sub>2</sub> models from within a mixed-fluid model can be used by adding 'Water' or  
904 'Carbon' to the model name (e.g., DixonCarbon; note that MagmaSat does not have  
905 this functionality).

906 Determination of the appropriate model to use with any sample is crucial to  
907 the correct application of these models, and so we stress the importance of under-  
908 standing how a model's calibration space relates to the sample at hand. VESIcal  
909 includes some built-in functionality for comparing melt compositions from user  
910 loaded data to those in the datasets upon which each of the VESIcal models is cal-  
911 ibrated using the method `calib_plot`. This can be visualized as a total alkalis vs  
912 silica (TAS) diagram (with fields and labels via the python `tasplot` library by J.  
913 Stevenson; <https://bitbucket.org/jsteven5/tasplot/src/master/>; Fig. 5a) or  
914 as any x-y plot in which x and y are oxides (Fig. 5b).

915 **Method structure:** `calib_plot(user_data=None, model='all', plot_type='`  
916 `TAS', zoom=None, save_fig=False)`

917 **Optional inputs:**

918 **user\_data**: The default value is None, in which case only the model calibration  
 919 set is plotted. User provided sample data describing the oxide composition  
 920 of one or more samples. Multiple samples can be passed as an BatchFile  
 921 object or pandas DataFrame. A single sample can be passed as a pandas  
 922 Series.

923 **model**: The default value is 'all', in which case all model calibration datasets  
 924 will be plotted. Otherwise, any model can be plotted by passing the name  
 925 of the model desired (e.g., 'Liu'). Multiple models can be plotted by passing  
 926 them as strings within a list (e.g., ['Liu', 'Dixon'])

927 **plot\_type**: The default value is 'TAS', which returns a total alkalis vs sil-  
 928 ica (TAS) diagram. Any two oxides can be plotted as an x-y plot by setting  
 929 plot\_type='xy' and specifying x- and y-axis oxides, e.g., x='SiO2', y='Al2O3'.

930 **zoom**: The default is None in which case axes will be set to the default of  
 931  $35 \leq x \leq 100$  wt% and  $0 \leq y \leq 25$  wt% for TAS type plots and the best values to  
 932 show the data for xy type plots. The user can pass "user\_data" to plot the  
 933 figure where the x and y axes are scaled down to zoom in and only show the  
 934 region surrounding the user\_data. A list of tuples may be passed to manually  
 935 specify x and y limits. Pass in data as [(x\_min, x\_max), (y\_min, y\_max)]. For  
 936 example, the default limits here would be passed in as [(35,100), (0,25)].

937 **save\_fig**: The default value is False, in which case the plot will be generated  
 938 and displayed but not saved. If the user wishes to save the figure, the desired  
 939 filename (including the file extension, e.g., .png) can be passed here. Note  
 940 that all plots in this Jupyter notebook can be saved by right clicking the plot  
 941 and choosing "Save Image As..."

## 942 **Outputs:**

943 A TAS or x-y plot of user data and model calibration data.

## 944 **Input**

```
945 v.calib_plot(user_data=myfile)
946 v.calib_plot(user_data=myfile, model='IaconoMarziano', plot_type='xy', x='
947 SiO2', y='K2O', save_fig=False)
948
```

## 950 **Output**





963 ditions. To calculate dissolved volatile concentrations thus requires computing the  
 964 equilibrium state of a system at fixed pressure and temperature over a range of bulk  
 965 volatile concentrations until a solution is found that satisfies the user defined fluid  
 966 composition.

967 First, the function makes an initial guess at the appropriate bulk volatile con-  
 968 centrations by finding the minimum dissolved volatile concentrations in the melt  
 969 at saturation, while asserting that the weight fraction of  $\text{H}_2\text{O}/(\text{H}_2\text{O}+\text{CO}_2)$  in the  
 970 system is equal to the user input mole fraction of  $\text{H}_2\text{O}/(\text{H}_2\text{O}+\text{CO}_2)$  in the fluid.  
 971 This is done by increasing the  $\text{H}_2\text{O}$  and  $\text{CO}_2$  concentrations appropriately until a  
 972 fluid phase is stable. Once fluid saturation is determined, the code then performs  
 973 directional, iterative, and progressively more refined searches, increasing the pro-  
 974 portion of  $\text{H}_2\text{O}$  or  $\text{CO}_2$  in the system if the mole fraction of  $\text{H}_2\text{O}$  calculated in the  
 975 fluid is greater than or less than that defined by the user, respectively. Four iterative  
 976 searches are performed; the precision of the match between the calculated and de-  
 977 fined  $\text{XH}_2\text{O}^{\text{fluid}}$  increases from 0.1 in the first iteration to 0.01, 0.001, and finally to  
 978 0.0001. Thus, the calculated dissolved volatile concentrations correspond to a system  
 979 with  $\text{XH}_2\text{O}^{\text{fluid}}$  within 0.0001 of the user defined value.

980 For non-MagmaSat models, dissolved volatile concentrations are calculated  
 981 directly from model equations.

## 982 Method structure:

983 Single sample: `calculate_dissolved_volatiles(sample, temperature,`  
 984 `pressure, X_fluid=1, verbose=False, model='MagmaSat').result`

985 BatchFile batch process: `myfile.calculate_dissolved_volatiles(  
 986 temperature, pressure, X_fluid=1, print_status=True, model='  
 987 MagmaSat')`

## 988 Standard inputs:

989 `sample, temperature, pressure, X_fluid, model` (see Section 3.0.1).

## 990 Unique optional inputs:

991 `verbose`: *Only for single sample calculations.* Default value is False in which  
 992 case  $\text{H}_2\text{O}$  and  $\text{CO}_2$  concentrations are returned. If set to True, additional  
 993 parameters are returned in a dictionary:  $\text{H}_2\text{O}$  and  $\text{CO}_2$  concentrations in the  
 994 fluid in mole fraction, temperature, pressure, and proportion of the fluid in  
 995 the system in wt%.

996 `print_status`: *Only for batch calculations.* The default value is True, in  
 997 which case the progress of the calculation will be printed to the terminal. The  
 998 user may desire to see the status of the calculation, as this particular function  
 999 can be quite slow, averaging between 3-5 seconds per sample.

## 1000 Calculated outputs:

1001 If the single-sample method is used, a dictionary with keys 'H2O' and 'CO2'  
 1002 corresponding to the calculated dissolved  $\text{H}_2\text{O}$  and  $\text{CO}_2$  concentrations in the  
 1003 melt is returned (plus additional variables 'temperature' in °C, 'pressure' in

1004 bars, 'XH2O.fl', 'XCO2.fl', and 'FluidProportion.wtper' (the proportion of  
1005 the fluid in the system in wt%) if `verbose` is set to `True`).

1006 If the `BatchFile` method is used, a pandas `DataFrame` is returned with sample  
1007 information plus calculated dissolved H<sub>2</sub>O and CO<sub>2</sub> concentrations in the  
1008 melt, the fluid composition in mole fraction, and the proportion of the fluid in  
1009 the system in wt%. Pressure (in bars) and Temperature (in °C) columns are  
1010 always returned.

### 1011 Input

```
1012 """Calculate dissolved volatiles for sample 10*"""
1013
1014 v.calculate_dissolved_volatiles(sample=sample_10, temperature=900.0, pressure
1015                               =2000.0, X_fluid=0.5, verbose=True).
1016                               result
```

### 1018 Output

```
1019 {'H2O.liq': 2.69352739399806,
1020  'CO2.liq': 0.0638439414375309,
1021  'XH2O.fl': 0.500092686493868,
1022  'XCO2.fl': 0.499907313506132,
1023  'FluidProportion.wt': 0.18407321260435108}
```

### 1026 Input

```
1027 """Calculate dissolved for all samples in an BatchFile object"""
1028
1029 dissolved = myfile.calculate_dissolved_volatiles(temperature=900.0, pressure=
1030                                                  2000.0, X_fluid=1, print_status=True)
1031
1032 dissolved
```

### 1033 Output

**Table 5.** Modeled dissolved volatile concentrations. *Note: This table has been truncated to display only the results of the calculation. The actual returned table would include all originally input user data in the leftmost columns followed by the calculation results. The complete table can be seen in the Jupyter notebook version of this manuscript.*

	User Input Data	H2O_liq_VESlcal	CO2_liq_VESlcal	Temperature_C_VESlcal	Pressure_bars_VESlcal	X_fluid_input_VESlcal	Model	Warnings
<b>Label</b>								
Kil3-6_1a	--	5.256561	0	900	2000	1	MagmaSat	
Kil3-6_3a	--	5.417720	0	900	2000	1	MagmaSat	
Kil3-6_4a	--	5.353421	0	900	2000	1	MagmaSat	
10*	--	4.984021	0	900	2000	1	MagmaSat	
19*	--	5.134419	0	900	2000	1	MagmaSat	
25	--	5.189068	0	900	2000	1	MagmaSat	
SAT-M12-1	--	5.810439	0	900	2000	1	MagmaSat	
SAT-M12-2	--	5.810439	0	900	2000	1	MagmaSat	
SAT-M12-4	--	5.810439	0	900	2000	1	MagmaSat	
samp. P1968a	--	6.484749	0	900	2000	1	MagmaSat	
samp. P1968b	--	6.473813	0	900	2000	1	MagmaSat	
samp. P1968c	--	6.482109	0	900	2000	1	MagmaSat	
samp. HPR3-1_XL-3	--	6.097630	0	900	2000	1	MagmaSat	
samp. HPR3-1_XL-4_INCL-1	--	6.138658	0	900	2000	1	MagmaSat	
AW-6	--	5.856636	0	900	2000	1	MagmaSat	
AW-46	--	5.879457	0	900	2000	1	MagmaSat	
KI-07	--	4.918430	0	900	2000	1	MagmaSat	

1034

### 3.4 Calculating equilibrium fluid compositions

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The `calculate_equilibrium_fluid_comp()` function calculates the composition of a fluid phase in equilibrium with a given silicate melt with known pressure, temperature, and dissolved H<sub>2</sub>O and CO<sub>2</sub> concentrations. The calculation is performed simply by calculating the equilibrium state of the given sample at the given conditions and determining if that melt is fluid saturated. If the melt is saturated, fluid composition and mass are reported back. If the calculation finds that the melt is not saturated at the given pressure and temperature, values of 0.0 will be returned for the H<sub>2</sub>O and CO<sub>2</sub> concentrations in the fluid.

1043

#### Method structure:

1044

1045

```
Single sample: calculate_equilibrium_fluid_comp(sample, temperature,
pressure, verbose=False, model='MagmaSat').result
```

1046

1047

```
BatchFile batch process: myfile.calculate_equilibrium_fluid_comp(
temperature, pressure=None, print_status=False, model='MagmaSat')
```

1048

#### Standard inputs:

1049

sample, temperature, pressure, model (see Section 3.0.1).

1050

#### Unique optional inputs:

1051

1052

1053

`verbose`: Only for single sample calculations. Default value is False, in which case H<sub>2</sub>O and CO<sub>2</sub> concentrations in the fluid in mol fraction are returned. If set to True, additional parameters are returned in a dictionary: H<sub>2</sub>O and

1054 CO<sub>2</sub> concentrations in the fluid, mass of the fluid in grams, and proportion of  
 1055 the fluid in the system in wt%.

1056 `print_status`: *Only for batch calculations.* The default value is False. If True  
 1057 is passed, the progress of the calculation will be printed to the terminal.

## 1058 **Calculated outputs:**

1059 If the single-sample method is used, a dictionary with keys 'H2O' and 'CO2'  
 1060 is returned (plus additional variables 'FluidMass\_grams' and 'FluidPropor-  
 1061 tion\_wtper' if `verbose` is set to True).

1062 If the BatchFile method is used, a pandas DataFrame is returned with sample  
 1063 information plus calculated equilibrium fluid compositions, mass of the fluid  
 1064 in grams, and proportion of the fluid in the system in wt%. Pressure (in bars)  
 1065 and Temperature (in °C) columns are always returned.

## 1066 **Input**

```
1067 """Calculate fluid composition for the extracted sample"""
1068 v.calculate_equilibrium_fluid_comp(sample=sample_10, temperature=900.0,
1069                                 pressure=100.0).result
1070
1071
```

## 1072 **Output**

```
1073 {'CO2': 0.00528661429366132, 'H2O': 0.994713385706339}
1074
1075
```

1076 Below we calculate equilibrium fluid compositions for all samples at a single  
 1077 temperature of 900 °C and a single pressure of 1,000 bars. Note that some samples  
 1078 in this dataset have quite low volatile concentrations (e.g., the Tucker et al. (2019)  
 1079 basalts from Kilauea), and so are below saturation at this P-T condition. The fluid  
 1080 composition for undersaturated samples is returned as values of 0 for both H<sub>2</sub>O and  
 1081 CO<sub>2</sub>.

## 1082 **Input**

```
1083 """Calculate fluid composition for all samples in an BatchFile object"""
1084 eqfluid = myfile.calculate_equilibrium_fluid_comp(temperature=900.0, pressure
1085                                                  =1000.0)
1086
1087 eqfluid
1088
```

**Table 6.** Isothermally modeled equilibrium fluid compositions. *Note: this table has been truncated to display only the results of the calculation. The actual returned table would include all originally input user data in the leftmost columns followed by the calculation results. The complete table can be seen in the Jupyter notebook version of this manuscript.*

Label	User Input Data	XH2O_fl_VESICAL	XCO2_fl_VESICAL	Temperature_C_VESICAL	Pressure_bars_VESICAL	Model	Warnings
Kii3-6_1a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions
Kii3-6_3a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions
Kii3-6_4a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions
10*	--	0.984531	0.015469	900	1000	MagmaSat	
19*	--	0.974997	0.025003	900	1000	MagmaSat	
25	--	0.990107	0.009893	900	1000	MagmaSat	
SAT-M12-1	--	1.000000	0.000000	900	1000	MagmaSat	
SAT-M12-2	--	1.000000	0.000000	900	1000	MagmaSat	
SAT-M12-4	--	1.000000	0.000000	900	1000	MagmaSat	
samp. P1968a	--	0.977773	0.022227	900	1000	MagmaSat	
samp. P1968b	--	0.996799	0.003201	900	1000	MagmaSat	
samp. P1968c	--	0.997028	0.002972	900	1000	MagmaSat	
samp. HPR3-1_XL-3	--	0.997770	0.002230	900	1000	MagmaSat	
samp. HPR3-1_XL-4_INCL-1	--	0.997273	0.002727	900	1000	MagmaSat	
AW-6	--	0.261572	0.738428	900	1000	MagmaSat	
AW-46	--	0.897441	0.102559	900	1000	MagmaSat	
KI-07	--	0.826014	0.173986	900	1000	MagmaSat	

Below, we calculate equilibrium fluid compositions for the same dataset using temperatures and pressures as defined in the input data (Table 3). Note that Samples “samp. HPR3-1\_XL-3” and “samp. HPR3-1\_XL-4\_INCL-1” have a user-defined value of 0.0 for temperature and pressure, respectively. VESICAL automatically skips the calculation of equilibrium fluids for these samples and returns a warning to the user, which are both printed to the terminal below and appended to the “Warnings” column in the returned data.

#### Input

```

1097 """Calculate fluid composition for all samples with unique pressure and
1098         temperature values for each sample.
1099 Pressure and temperature values are taken from columns named "Press" and "
1100         Temp" in the example BatchFile"""
1101
1102 eqfluid_wtemps = myfile.calculate_equilibrium_fluid_comp(temperature='Temp',
1103         pressure='Press')
1104
1105 eqfluid_wtemps
1106

```

#### Output

```

1108 UserWarning: Temperature for sample samp. HPR3-1_XL-3 is <=0. Skipping sample.
1109 UserWarning: Pressure for sample samp. HPR3-1_XL-4_INCL-1 is <=0. Skipping sample.
1110

```

**Table 7.** Modeled equilibrium fluid compositions with unique temperatures. Warnings “Bad temperature” and “Bad pressure” indicate that no data (or 0.0 value data) was given for the temperature or pressure of that sample, in which case the calculation of that sample is skipped. *Note: this table has been truncated to display only the results of the calculation. The actual returned table would include all originally input user data in the leftmost columns followed by the calculation results. The complete table can be seen in the Jupyter notebook version of this manuscript.*

	User Input Data	XH2O_fl_VESICAL	XCO2_fl_VESICAL	Model	Warnings
Label					
Kil3-6_1a	--	0.586164	0.413836	MagmaSat	
Kil3-6_3a	--	0.286160	0.713840	MagmaSat	
Kil3-6_4a	--	0.377439	0.622561	MagmaSat	
10*	--	0.892371	0.107629	MagmaSat	
19*	--	0.918888	0.081112	MagmaSat	
25	--	0.955803	0.044197	MagmaSat	
SAT-M12-1	--	1.000000	0.000000	MagmaSat	
SAT-M12-2	--	1.000000	0.000000	MagmaSat	
SAT-M12-4	--	1.000000	0.000000	MagmaSat	
samp. P1968a	--	0.998764	0.001236	MagmaSat	
samp. P1968b	--	0.998686	0.001314	MagmaSat	
samp. P1968c	--	0.998831	0.001169	MagmaSat	
samp. HPR3-1_XL-3	--	NaN	NaN	MagmaSat	Calculation skipped. Bad temperature.
samp. HPR3-1_XL-4_INCL-1	--	NaN	NaN	MagmaSat	Calculation skipped. Bad pressure.
AW-6	--	0.000000	0.000000	MagmaSat	Sample not saturated at these conditions
AW-46	--	0.492213	0.507787	MagmaSat	
KI-07	--	0.681758	0.318242	MagmaSat	

### 3.4.1 Converting fluid composition units

The fluid composition is always returned in units of mol fraction. Two functions exist to transform only the H<sub>2</sub>O-CO<sub>2</sub> fluid composition between mol fraction and wt% and can be applied to returned data sets from calculations. Both functions require that the user provide the dataframe containing fluid composition information plus the names of the columns corresponding to the H<sub>2</sub>O and CO<sub>2</sub> concentrations in the fluid. The default values for column names are set to those that may be returned by VESICAL core calculations, such that they need not be specified unless the user has changed them or is supplying their own data (e.g., imported data not processed through a core calculation).

#### Method structure:

```
Mol fraction to wt%: fluid_molfrac_to_wt(data, H2O_colname='
XH2O_fl_VESICAL', CO2_colname='XCO2_fl_VESICAL')
```

```
Wt% to mol fraction: fluid_wt_to_molfrac(data, H2O_colname='
H2O_fl_wt', CO2_colname='CO2_fl_wt')
```

#### Required inputs:

1128 data: A pandas DataFrame containing columns for H<sub>2</sub>O and CO<sub>2</sub> concentra-  
 1129 tions in the fluid.

1130 **Optional inputs:**

1131 H2O\_colname and CO2\_colname: The default values are 'XH2O\_fl' and  
 1132 'XCO2\_fl' if input data are in mol fraction or 'H2O\_fl\_wt' and 'CO2\_fl\_wt'  
 1133 if the data are in wt%. Strings containing the name of the columns corre-  
 1134 sponding to the H<sub>2</sub>O and CO<sub>2</sub> concentrations in the fluid.

1135 **Calculated outputs:**

1136 The original data passed plus newly calculated values are returned in a  
 1137 DataFrame.

1138 **Input**

```
1139 """Converting from mol fraction to wt%"""
1140 eqfluid_wt = v.fluid_molfrac_to_wt(eqfluid)
1141 eqfluid_wt
1142
1143
```

**Table 8.** Equilibrium fluid compositions converted from mol fraction to wt%. *Note: this table has been truncated to display only the results of the calculation. The actual returned table would include all originally input user data in the leftmost columns followed by the calculation results. The complete table can be seen in the Jupyter notebook version of this manuscript.*

	User Input Data	XH2O_fl_VESical	XCO2_fl_VESical	Temperature_C_VESical	Pressure_bars_VESical	Model	Warnings	H2O_fl_wt	CO2_fl_wt
<b>Label</b>									
Kil3-6_1a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions		
Kil3-6_3a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions		
Kil3-6_4a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions		
10*	--	0.984531	0.015469	900	1000	MagmaSat		96.3044	3.69556
19*	--	0.974997	0.025003	900	1000	MagmaSat		94.1062	5.89383
25	--	0.990107	0.009893	900	1000	MagmaSat		97.6179	2.38209
SAT-M12-1	--	1.000000	0.000000	900	1000	MagmaSat		100	0
SAT-M12-2	--	1.000000	0.000000	900	1000	MagmaSat		100	0
SAT-M12-4	--	1.000000	0.000000	900	1000	MagmaSat		100	0
samp. P1968a	--	0.977773	0.022227	900	1000	MagmaSat		94.7402	5.25979
samp. P1968b	--	0.996799	0.003201	900	1000	MagmaSat		99.2217	0.778256
samp. P1968c	--	0.997028	0.002972	900	1000	MagmaSat		99.2773	0.722709
samp. HPR3-1_XL-3	--	0.997770	0.002230	900	1000	MagmaSat		99.457	0.542973
samp. HPR3-1_XL-4_INCL-1	--	0.997273	0.002727	900	1000	MagmaSat		99.3367	0.6633
AW-6	--	0.261572	0.738428	900	1000	MagmaSat		12.6667	87.3333
AW-46	--	0.897441	0.102559	900	1000	MagmaSat		78.1798	21.8202
KI-07	--	0.826014	0.173986	900	1000	MagmaSat		66.0315	33.9685

1144 **Input**



1145  
1146  
1147  
1148

```

"""Converting from wt% to mol fraction"""
eqfluid_mol = v.fluid_wt_to_molfrac(eqfluid_wt)
eqfluid_mol

```

**Table 9.** Equilibrium fluid compositions converted from wt% to mol fraction. *Note: this table has been truncated to display only the results of the calculation. The actual returned table would include all originally input user data in the leftmost columns followed by the calculation results. The complete table can be seen in the Jupyter notebook version of this manuscript.*

Label	User Input Data	XH2O_fl _VESical	XCO2_fl _VESical	Temperature_C _VESical	Pressure_bars _VESical	Model	Warnings	H2O_fl_wt	CO2_fl_wt	XH2O_fl	XCO2_fl
Kil3-6_1a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions				
Kil3-6_3a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions				
Kil3-6_4a	--	0.000000	0.000000	900	1000	MagmaSat	Sample not saturated at these conditions				
10*	--	0.984531	0.015469	900	1000	MagmaSat		96.3044	3.69556	0.984531	0.0154691
19*	--	0.974997	0.025003	900	1000	MagmaSat		94.1062	5.89383	0.974997	0.0250027
25	--	0.990107	0.009893	900	1000	MagmaSat		97.6179	2.38209	0.990107	0.0098927
SAT-M12-1	--	1.000000	0.000000	900	1000	MagmaSat		100	0	1	0
SAT-M12-2	--	1.000000	0.000000	900	1000	MagmaSat		100	0	1	0
SAT-M12-4	--	1.000000	0.000000	900	1000	MagmaSat		100	0	1	0
samp. P1968a	--	0.977773	0.022227	900	1000	MagmaSat		94.7402	5.25979	0.977773	0.0222267
samp. P1968b	--	0.996799	0.003201	900	1000	MagmaSat		99.2217	0.778256	0.996799	0.0032013
samp. P1968c	--	0.997028	0.002972	900	1000	MagmaSat		99.2773	0.722709	0.997028	0.00297183
samp. HPR3-1_XL-3	--	0.997770	0.002230	900	1000	MagmaSat		99.457	0.542973	0.99777	0.00223037
samp. HPR3-1_XL-4_INCL-1	--	0.997273	0.002727	900	1000	MagmaSat		99.3367	0.6633	0.997273	0.00272658
AW-6	--	0.261572	0.738428	900	1000	MagmaSat		12.6667	87.3333	0.261572	0.738428
AW-46	--	0.897441	0.102559	900	1000	MagmaSat		78.1798	21.8202	0.897441	0.102559
KI-07	--	0.826014	0.173986	900	1000	MagmaSat		66.0315	33.9685	0.826014	0.173986

1150

### 3.5 Calculating saturation pressures

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1155  
1156

The `calculate_saturation_pressure()` function calculates the minimum pressure at which a given silicate melt with known temperature and H<sub>2</sub>O and CO<sub>2</sub> concentrations would be saturated with fluid. For MagmaSat, this is calculated by finding the pressure at which the smallest amount of vapor is present. This function also calculates the composition of the vapor in equilibrium with the melt at those conditions.

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1161

The function works by calculating the equilibrium state of the given melt at very high pressure (20,000 bars) and then decreasing the pressure in steps of 1,000 bars until the mass of vapor is >0 grams. At this point, the pressure space is narrowed and searched in steps of 100 bars and then in steps of 10 bars until the saturation pressure is found. Thus, these calculations are accurate to 10 bars.

1162  
1163  
1164  
1165

For non-MagmaSat models, we use Brent's minimization method (via `scipy's` `root_scalar` optimization function) to find the pressure that satisfies the computational constraints. This is achieved by iterative calculation of the dissolved volatile concentration over a range of pressures and minimizing the difference between com-

1166 puted and given concentrations. This is only practical for non-MagmaSat models,  
 1167 where the dissolved volatiles calculation is extremely fast.

### 1168 Method structure:

1169 Single sample: `calculate_saturation_pressure(sample, temperature,`  
 1170 `verbose=False, model='MagmaSat').result`

1171 BatchFile batch process: `myfile.calculate_saturation_pressure(  
 1172 temperature, print_status=True, model='MagmaSat')`

### 1173 Standard inputs:

1174 `sample, temperature, model` (see Section 3.0.1).

### 1175 Unique optional inputs:

1176 `verbose`: *Only for single sample calculations.* Default value is False in which  
 1177 case the saturation pressure in bars is returned. If set to True, additional  
 1178 parameters are returned in a dictionary: saturation pressure in bars, H<sub>2</sub>O and  
 1179 CO<sub>2</sub> concentrations in the fluid, mass of the fluid in grams, and proportion of  
 1180 the fluid in the system in wt%.

1181 `print_status`: *Only for batch calculations.* The default value is True, in  
 1182 which case the progress of the calculation will be printed to the terminal.

### 1183 Calculated outputs:

1184 If the single-sample method is used, the saturation pressure in bars is re-  
 1185 turned as a numerical value (float) (plus additional variables 'XH2O\_fl',  
 1186 'XCO2\_fl', 'FluidMass\_grams', and 'FluidProportion\_wtper' if `verbose` is set  
 1187 to True).

1188 If the BatchFile method is used, a pandas DataFrame is returned with sample  
 1189 information plus calculated saturation pressures, equilibrium fluid composi-  
 1190 tions, mass of the fluid in grams, and proportion of the fluid in the system in  
 1191 wt%. Temperature (in °C) is always returned.

### 1192 Input

```
1193 """Calculate the saturation pressure of the single sample we defined in
1194                               Section 3.1.2 at 925 degrees C"""
1195 v.calculate_saturation_pressure(sample=mysample, temperature=925.0, verbose=
1196                               True).result
1197
1198
```

### 1199 Output

```
1200 {'SaturationP_bars': 2720,
1201  'FluidMass_grams': 0.0016655984224872,
1202  'FluidProportion_wt': 0.0015635017577088073,
1203  'XH2O_fl': 0.825802671679744,
1204  'XCO2_fl': 0.174197328320256}
1205
1206
```

1207 Input

1208

1209

1210

1211

1212

1213

```

"""Calculate the saturation pressure for all samples in an BatchFile object
    at 925 degrees C"""
satPs = myfile.calculate_saturation_pressure(temperature=925.0)
satPs

```

**Table 10.** Isothermally modeled saturation pressures. *Note: this table has been truncated to display only the results of the calculation. The actual returned table would include all originally input user data in the leftmost columns followed by the calculation results. The complete table can be seen in the Jupyter notebook version of this manuscript.*

Label	User Input Data	SaturationP_bars _VESical	Temperature_C _VESical	XH2O_fl _VESical	XCO2_fl _VESical	FluidMass_grams _VESical	FluidSystem_wt _VESical	Model	Warnings
Kii3-6_1a	--	80	925	0.439592	0.560408	0.000026	0.000029	MagmaSat	
Kii3-6_3a	--	140	925	0.248725	0.751275	0.000620	0.000696	MagmaSat	
Kii3-6_4a	--	120	925	0.303412	0.696588	0.000332	0.000372	MagmaSat	
10*	--	2370	925	0.785443	0.214557	0.001277	0.001221	MagmaSat	
19*	--	3430	925	0.685666	0.314334	0.000226	0.000215	MagmaSat	
25	--	2580	925	0.825858	0.174142	0.000899	0.000854	MagmaSat	
SAT-M12-1	--	540	925	1.000000	0.000000	0.005197	0.005125	MagmaSat	
SAT-M12-2	--	1490	925	1.000000	0.000000	0.015744	0.015163	MagmaSat	
SAT-M12-4	--	2340	925	1.000000	0.000000	0.000981	0.000929	MagmaSat	
samp. P1968a	--	1030	925	0.970939	0.029061	0.005539	0.005308	MagmaSat	
samp. P1968b	--	1640	925	0.970540	0.029460	0.003193	0.003016	MagmaSat	
samp. P1968c	--	1590	925	0.972952	0.027048	0.000828	0.000783	MagmaSat	
samp. HPR3-1_XL-3	--	1920	925	0.946905	0.053095	0.001085	0.001026	MagmaSat	
samp. HPR3-1_XL-4_INCL-1	--	1600	925	0.947015	0.052985	0.002230	0.002120	MagmaSat	
AW-6	--	1200	925	0.229039	0.770961	0.000322	0.000317	MagmaSat	
AW-46	--	4640	925	0.439714	0.560286	0.000411	0.000391	MagmaSat	
KI-07	--	1450	925	0.678808	0.321192	0.002972	0.002886	MagmaSat	

1214 Input

1215

1216

1217

1218

1219

1220

1221

```

"""Calculate the saturation pressure for all samples in an BatchFile object,
    taking temperature
    values from a column named "Temp" in the BatchFile"""
satPs_wtemps = myfile.calculate_saturation_pressure(temperature="Temp")
satPs_wtemps

```

**Table 11.** Modeled saturation pressures with unique temperatures. The warning “Bad temperature” indicates that no data (or 0.0 value data) was given for the temperature of that sample, in which case the calculation of that sample is skipped. *Note: this table has been truncated to display only the results of the calculation. The actual returned table would include all originally input user data in the leftmost columns followed by the calculation results. The complete table can be seen in the Jupyter notebook version of this manuscript.*

Label	User Input Data	SaturationP_bars_VESical	XH2O_fl_VESical	XCO2_fl_VESical	FluidMass_grams_VESical	FluidSystem_wt_VESical	Model	Warnings
Kil3-6_1a	--	70	0.525553	0.474447	0.000710314	0.000797025	MagmaSat	
Kil3-6_3a	--	130	0.281991	0.718009	0.000449053	0.00050435	MagmaSat	
Kil3-6_4a	--	110	0.344636	0.655364	0.000370296	0.000415397	MagmaSat	
10*	--	2400	0.807034	0.192966	0.00029124	0.00027857	MagmaSat	
19*	--	3440	0.710745	0.289255	0.000584324	0.000555328	MagmaSat	
25	--	2660	0.845161	0.154839	0.000211825	0.000201233	MagmaSat	
SAT-M12-1	--	560	1	0	0.0229218	0.0226009	MagmaSat	
SAT-M12-2	--	1560	1	0	0.00331538	0.00319308	MagmaSat	
SAT-M12-4	--	2390	1	0	0.0186879	0.0177036	MagmaSat	
samp. P1968a	--	1020	0.971529	0.0284714	0.00694889	0.00666009	MagmaSat	
samp. P1968b	--	1630	0.970816	0.0291835	0.00258244	0.00243943	MagmaSat	
samp. P1968c	--	1580	0.973228	0.0267722	0.000412392	0.000389913	MagmaSat	
samp. HPR3-1_XL-3	--						MagmaSat	Calculation skipped. Bad temperature.
samp. HPR3-1_XL-4_INCL-1	--	1600	0.947004	0.0529965	0.00234923	0.0022335	MagmaSat	
AW-6	--	1270	0.224501	0.775499	4.1948e-05	4.13078e-05	MagmaSat	
AW-46	--	4740	0.441496	0.558504	0.000345244	0.00032851	MagmaSat	
KI-07	--	1530	0.673211	0.326789	0.0032638	0.00316864	MagmaSat	

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### 3.6 Calculating isobars and isopleths

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In this example, we demonstrate how isobars (lines of constant pressure) and isopleths (lines of constant fluid composition) can be calculated for any one composition. A single melt composition can be extracted from a loaded batch file, or a composition can be entered by hand and stored within a dictionary. Due to computational intensity, isobars and isopleths can only be computed for one sample composition at a time.

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Once a single composition is defined, conditions over which to calculate isobars and isopleths must be specified. The generated plot is isothermal, so only one temperature can be chosen. Isobars and isopleths can be calculated for any number of pressures or  $\text{XH}_2\text{O}^{fluid}$  values, respectively, passed as lists.

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The calculation is performed by iterating through possible concentrations of  $\text{H}_2\text{O}$  and  $\text{CO}_2$  and calculating the equilibrium state for the system. The iteration begins at a fixed  $\text{H}_2\text{O}$  concentration, increasing the  $\text{CO}_2$  concentration in steps of 0.1 wt% until a fluid phase is stable. The  $\text{H}_2\text{O}$  concentration is then increased by 0.5 wt% and  $\text{CO}_2$  is again increased from 0 until a fluid phase is stable. This process is repeated for  $\text{H}_2\text{O}$  values ranging from 0–15 wt%. The  $\text{H}_2\text{O}$  and  $\text{CO}_2$  concentrations from each system for which a fluid phase was found to be stable are saved and written to a pandas DataFrame, which is returned upon completion of the calculation.

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Isobars and isopleths are computed at fixed  $\text{H}_2\text{O}$ - $\text{CO}_2$  points for any given pressure. To generate curves using the MagmaSat model, polynomials are fit to computed points using numpy’s polyfit method. This can be optionally disabled by setting `smooth_isobars` or `smooth_isopleths` to False. The curvature of the isobars depends strongly on the number of points used to fit a polynomial, deemed

1246 “control points”, with curve fits becoming more accurate to the model as the num-  
 1247 ber of control points increases. We found that above five control points, changes  
 1248 to the shape of the curve fits becomes negligible. Thus, as a compromise between  
 1249 accuracy and computation time, and to maintain consistency, MagmaSat isobars are  
 1250 always computed with 5 control points at  $\text{XH}_2\text{O}^{fluid}$  values of 0, 0.25, 0.5, 0.75, and  
 1251 1. Because non-MagmaSat models compute extremely quickly, all non-MagmaSat  
 1252 models use 51 control points per isobar and do not utilize polynomial fits to the data  
 1253 by default.

#### 1254 **Method structure:**

```
1255     Only single sample calculations. calculate_isobars_and_isopleths(sample
1256     , temperature, pressure_list, isopleth_list=None, smooth_isobars=
1257     True, smooth_isopleths=True, print_status=True, model="MagmaSat").
1258     result
```

#### 1259 **Standard inputs:**

1260 `sample, temperature, model` (see Section 3.0.1).

#### 1261 **Unique required inputs:**

1262 `pressure_list`: A list of all pressures in bars at which to calculate isobars. If  
 1263 only one value is passed it can be as float instead of list.

#### 1264 **Unique optional inputs:**

1265 `isopleth_list`: The default value is None in which case only isobars will  
 1266 be calculated. A list of all fluid composition values, in mole fraction  $\text{H}_2\text{O}$   
 1267 ( $\text{XH}_2\text{O}^{fluid}$ ), at which to calculate isopleths. Values can range from 0–1. If  
 1268 only one value is passed it can be as float instead of list. N.b. that, due to  
 1269 the method of isobar smoothing using control points as outlined above, each  
 1270 isopleth value passed here not equal to one of the five standard control point  
 1271 values (0, 0.25, 0.5, 0.75, or 1) will result in an an additional control point  
 1272 being used to smooth the isobars. Thus, entering additional isopleth values  
 1273 results not only in more isopleth outputs but also in “smoother” (i.e., more  
 1274 well constrained) isobars.

1275 `smooth_isobars` and `smooth_isopleths`: The default value for both of these  
 1276 arguments is True, in which case polynomials will be fit to the computed data  
 1277 points.

1278 `print_status`: The default value is True. If True, the progress of the calcula-  
 1279 tions will be printed to the terminal.

#### 1280 **Calculated outputs:**

1281 The function returns two pandas DataFrames: the first has isobar data, and  
 1282 the second has isopleth data. Columns in the isobar dataframe are ‘Pressure’,  
 1283 ‘H2Omelt’, and ‘CO2melt’, corresponding to pressure in bars and dissolved  
 1284  $\text{H}_2\text{O}$  and  $\text{CO}_2$  in the melt in wt%. Columns in the isopleth dataframe are

1285 'XH2O.fl', 'H2O.liq', and 'CO2.liq', corresponding to  $\text{XH}_2\text{O}^{fluid}$  and dis-  
 1286 solved  $\text{H}_2\text{O}$  and  $\text{CO}_2$  in the melt in wt%.

1287 Input

```
1288 """Define all variables to be passed to the function for calculating isobars
1289                                     and isopleths"""
1290
1291 """Define the temperature in degrees C"""
1292 temperature = 1200.0
1293
1294 """Define a list of pressures in bars:"""
1295 pressures = [1000.0, 2000.0, 3000.0]
1296
```

1297 Next, the  $\text{H}_2\text{O}$  and  $\text{CO}_2$  dissolved in the melt at saturation is calculated at the  
 1298 specified temperature and over the range of specified pressures. Note that, because  
 1299 this function calculates two things (isobars and isopleths), two variable names must  
 1300 be given (below, "isobars, isopleths"). This calculation can be quite slow, and so it  
 1301 is recommended to set `print_status` to `True`.

1302 Input

```
1303 isobars, isopleths = v.calculate_isobars_and_isopleths(sample=sample_10,
1304                                                       temperature=temperature, pressure_list
1305                                                       =pressures, isopleth_list=[0.25,0.5,0.
1306                                                       75]).result
1307
1308
```

1309 Output

```
1310 Calculating isobar at 1000.0 bars
1311 done.
1312 Calculating isobar at 2000.0 bars
1313 done.
1314 Calculating isobar at 3000.0 bars
1315 done.
1316 Done!
1317
1318
```

### 1319 3.7 Calculating degassing paths

1320 A degassing path is a series of volatile concentrations both in the melt and  
 1321 fluid that a magma will follow during decompression. In the calculation, the saturation  
 1322 pressure is computed, and then the system is equilibrated along a trajectory of  
 1323 decreasing pressure values at discrete steps. The default number of steps to calcu-  
 1324 late is 50, but this can be defined by the user by setting the argument `steps` to any  
 1325 integer value. A detailed explanation of how non-MagmaSat models handle the cal-  
 1326 culation of mixed-fluid composition can be found in the supplement (Supplementary  
 1327 Text S2). If so desired, this calculation can be performed for any initial pressure,  
 1328 but the default is the saturation pressure. If a pressure is specified that is above the  
 1329 saturation pressure, the calculation will simply proceed from the saturation pressure,  
 1330 since the magma cannot degas until it reaches saturation.

1331 Completely open-system, completely closed-system or partially open-system  
 1332 degassing paths can be calculated by specifying what proportion of the fluid to frac-  
 1333 tionate. The fluid fractionation value can range between 0 (closed-system: no fluid  
 1334 is removed, all is retained at each pressure step) and 1 (open-system: all fluid is  
 1335 removed, none is retained at each pressure step). Closed and partially open-system

1336 runs allow the user to specify the initial presence of exsolved fluid that is in equilib-  
 1337 rium with the melt at the starting pressure.

### 1338 **Method structure:**

```
1339     Only single-sample calculations. calculate_degassing_path(sample,
1340         temperature, pressure='saturation', fractionate_vapor=0.0,
1341         init_vapor=0.0, steps=50, model='MagmaSat').result
```

### 1342 **Standard inputs:**

1343 sample, temperature, model (see Section 3.0.1).

### 1344 **Unique optional inputs:**

1345 **pressure:** The pressure at which to begin the degassing calculations, in  
 1346 bars. Default value is 'saturation', which runs the calculation with the initial  
 1347 pressure at the saturation pressure. If a pressure greater than the saturation  
 1348 pressure is input, the calculation will start at saturation, since this is the first  
 1349 pressure at which any degassing will occur.

1350 **fractionate\_vapor:** Proportion of vapor removed at each pressure step.  
 1351 Default value is 0.0 (completely closed-system degassing). Specifies the type  
 1352 of calculation performed, either closed system (0.0) or open system (1.0)  
 1353 degassing. If any value between <1.0 is chosen, user can also specify the  
 1354 'init\_vapor' argument (see below). A value in between 0 and 1 will remove  
 1355 that proportion of vapor at each step. For example, for a value of 0.2, the  
 1356 calculation will remove 20% of the vapor and retain 80% of the vapor at each  
 1357 pressure step.

1358 **init\_vapor:** Default value is 0.0. Specifies the amount of vapor (in wt%)  
 1359 coexisting with the melt before degassing.

1360 **steps:** Default value is 50. Specifies the number of steps in pressure space at  
 1361 which to calculate dissolved volatile concentrations.

### 1362 **Calculated outputs:**

1363 The function returns a pandas DataFrame with columns as: 'Pressure\_bars',  
 1364 'H2O\_liq' and 'CO2\_liq' (the concentration of H<sub>2</sub>O and CO<sub>2</sub> in the melt,  
 1365 in wt%), 'XH2O\_fl' and 'XCO2\_fl' (the composition of the H<sub>2</sub>O-CO<sub>2</sub> fluid,  
 1366 in mol fraction), and 'FluidProportion\_wt' (the proportion of fluid in the  
 1367 fluid-melt system, in wt%).

### 1368 **Input**

```
1369     temp = 1200 #temperature in degrees C
1370
1371     """Calculate open, closed, and closed + 2 wt% initial vapor"""
1372     closed_df = v.calculate_degassing_path(sample=sample_10, temperature=temp).
1373         result
1374
```

```

1375 open_df = v.calculate_degassing_path(sample=sample_10, temperature=temp,
1376                                     fractionate_vapor=1.0).result
1377 half_df = v.calculate_degassing_path(sample=sample_10, temperature=temp,
1378                                     fractionate_vapor=0.5).result
1379 exsolved_df = v.calculate_degassing_path(sample=sample_10, temperature=temp,
1380                                           init_vapor=2.0).result
1381
1382 """Calculate closed-system degassing starting from a pressure of 2000 bars"""
1383 start2000_df = v.calculate_degassing_path(sample=sample_10, temperature=temp,
1384                                           pressure=2000.0).result
1385

```

### 1386 3.8 Plotting

1387 After calculating isobars, isopleths, and degassing paths, any or all of these  
1388 may be plotted in an H<sub>2</sub>O versus CO<sub>2</sub> plot with one simple function call. The  
1389 plot will be printed directly in the notebook or, if the code is run as script in a  
1390 command line, the plot will appear in its own window, at which point it can be  
1391 saved as an image file. VESICAL's `plot` function takes in lists of pandas DataFrames  
1392 with calculated isobar, isopleth, and degassing path information (e.g., output from  
1393 `calculate_isobars_and_isopleths` or `calculate_degassing_path()`) and plots  
1394 data as isobars (lines of constant pressure), isopleths (lines of constant fluid compo-  
1395 sition), and degassing paths (lines indicating the concentrations of H<sub>2</sub>O and CO<sub>2</sub> in  
1396 a melt equilibrated along a path of decreasing pressure).

1397 Labels can be assigned to isobars, isopleths, and/or degassing paths separately.  
1398 Any or all of these data can be passed to the `plot` function. Multiple sets of plot-  
1399 table data can be passed. For example, isobars calculated with two different models  
1400 can be passed to the `isobars` argument as a list.

1401 VESICAL's plotting function is entirely based on python's matplotlib li-  
1402 brary, which comes standard with many installations of python. With matplotlib,  
1403 users can create a large variety of plots (note that direct matplotlib function-  
1404 ality is used to create custom plots in several of this manuscript's supplement-  
1405 ary Jupyter notebooks), and users should refer to the matplotlib documentation  
1406 (<https://matplotlib.org/3.2.1/index.html>) if more complex plotting is desired.  
1407 If preferred, VESICAL outputs can be saved to an Excel or CSV file (see Section  
1408 3.10), and plotting can be done in any plotting program desired (e.g., MS Excel).

1409 The function returns both fig and axes matplotlib objects, which can be fur-  
1410 ther edited by the user or plotted directly. Following matplotlib convention, the  
1411 results of `plot()` should be saved to objects such as fig, ax as:

```
1412 fig, ax = v.plot([options])
```

1413 where [options] represents any optional inputs as defined here. Variables fig and ax  
1414 can then be edited further using matplotlib tools. For example, the user might wish  
1415 to set the minimum x-axis value to 0.5 as:

```
1416 ax.set_xlim(left=0.5)
```

1417 In Jupyter Notebook, a plot is automatically shown, but in the command line, the  
1418 plot will only display after executing `v.show()`.

### 1419 Method structure:



```

1420 plot(isobars=None, isopleths=None, degassing_paths=None, custom_H2O
1421      =None, custom_CO2=None, isobar_labels=None, isopleth_labels=None
1422      , degassing_path_labels=None, custom_labels=None, custom_colors=
1423      "VESIcal", custom_symbols=None, markersize=10, save_fig=False,
1424      extend_isobars_to_zero=True, smooth_isobars=False, smooth_isopleths
1425      =False)

```

## 1426 Optional inputs:

1427 **isobars:** DataFrame object containing isobar information as calculated by  
 1428 `calculate_isobars_and_isopleths()`. Or a list of DataFrame objects.

1429 **isopleths:** DataFrame object containing isopleth information as calculated  
 1430 by `calculate_isobars_and_isopleths()`. Or a list of DataFrame objects.

1431 **degassing\_paths:** List of DataFrames with degassing information as gener-  
 1432 ated by `calculate_degassing_path()`.

1433 **custom\_H2O:** List of floats or array-like shapes of H<sub>2</sub>O concentration values  
 1434 to plot as points. For example `myfile.get_data()['H2O']` is one array-like  
 1435 shape (here, `pandas.Series`) of H<sub>2</sub>O values. Must be passed with `custom_CO2`  
 1436 and must be same length as `custom_CO2`.

1437 **custom\_CO2:** List of floats or array-like shapes of CO<sub>2</sub> values to plot as  
 1438 points. For example `myfile.get_data()['CO2']` is one array-like shape of  
 1439 CO<sub>2</sub> values. Must be passed with `custom_H2O` and must be same length as  
 1440 `custom_H2O`.

1441 **isobar\_labels:** Labels for the plot legend. Default is `None`, in which case  
 1442 each plotted line will be given the generic legend name of “Isobars n”, with  
 1443 n referring to the nth isobars passed. Isobar pressure is given in parentheses.  
 1444 The user can pass their own labels as a list of strings. If more than one set  
 1445 of isobars is passed, the labels should refer to each set of isobars, not each  
 1446 pressure.

1447 **isopleth\_labels:** Labels for the plot legend. Default is `None`, in which case  
 1448 each plotted isopleth will be given the generic legend name of “Isopleth n”,  
 1449 with n referring to the nth isopleths passed. Isopleth XH<sub>2</sub>O values are given  
 1450 in parentheses. The user can pass their own labels as a list of strings. If  
 1451 more than one set of isopleths is passed, the labels should refer to each set of  
 1452 isopleths, not each XH<sub>2</sub>O value.

1453 **degassing\_path\_labels:** Labels for the plot legend. Default is `None`,  
 1454 in which case each plotted line will be given the generic legend name of  
 1455 “Pathn”, with n referring to the nth degassing path passed. The user can  
 1456 pass their own labels as a list of strings.

1457 **custom\_labels:** Labels for the plot legend. Default is `None`, in which case  
 1458 each group of custom points will be given the generic legend name of “Cus-

1459 tomn”, with n referring to the nth degassing path passed. The user can pass  
 1460 their own labels as a list of strings.

1461 **custom\_colors** and **custom\_symbols**: Custom colors and symbol shapes can  
 1462 be specified for (**custom\_H2O**, **custom\_CO2**) points. A list of color values or  
 1463 symbol types readable by Matplotlib (see Matplotlib documentation) can be  
 1464 entered. The length of this list must be equal to the lengths of **custom\_H2O**  
 1465 and **custom\_CO2**. If nothing is specified for **custom\_colors**, VESICAL’s default  
 1466 colors will be used. If nothing is specified for **custom\_symbols**, all points will  
 1467 be plotted as filled circles.

1468 **markersize**: The size of the symbols can be specified here. If not specified,  
 1469 the default value is marker size 10.

1470 **save\_fig**: Default value is False, in which case the figure will not be saved.  
 1471 If a string is passed, the figure will be saved with the string as the filename.  
 1472 The string must include the file extension.

1473 **Advanced inputs**: Most users will not need to use these inputs.

1474 **extend\_isobars\_to\_zero**: If set to True (the default), isobars will be ex-  
 1475 tended to the plot axes, which are at x=0 and y=0, even if there is a finite  
 1476 solubility at zero partial pressure.

1477 **smooth\_isobars** and **smooth\_isopleths**: If set to True, isobar or iso-  
 1478 pleth data will be fit to a polynomial and plotted. If set to False (the  
 1479 default), the raw input data will be plotted. Note that MagmaSat  
 1480 **calculate\_isobars\_and\_isopleths()** calculations return already  
 1481 “smoothed” data (that is, the raw data are fit to polynomials before be-  
 1482 ing returned). Raw “unsmoothed” data can be returned by MagmaSat  
 1483 **calculate\_isobars\_and\_isopleths()** (see documentation on this method).

1484 **Calculated outputs**:

1485 The function returns fig and axes matplotlib objects defining a plot with  
 1486 x-axis as H<sub>2</sub>O wt% in the melt and y-axis as CO<sub>2</sub> wt% in the melt. Isobars,  
 1487 or lines of constant pressure at which the sample magma composition is sat-  
 1488 urated, and isopleths, or lines of constant fluid composition at which the  
 1489 sample magma composition is saturated, are plotted if passed. Degassing  
 1490 paths, or the concentration of dissolved H<sub>2</sub>O and CO<sub>2</sub> in a melt equilibrated  
 1491 along a path of decreasing pressure, is plotted if passed.

### 1492 **3.8.1 A simple example: Isobars and isopleths**

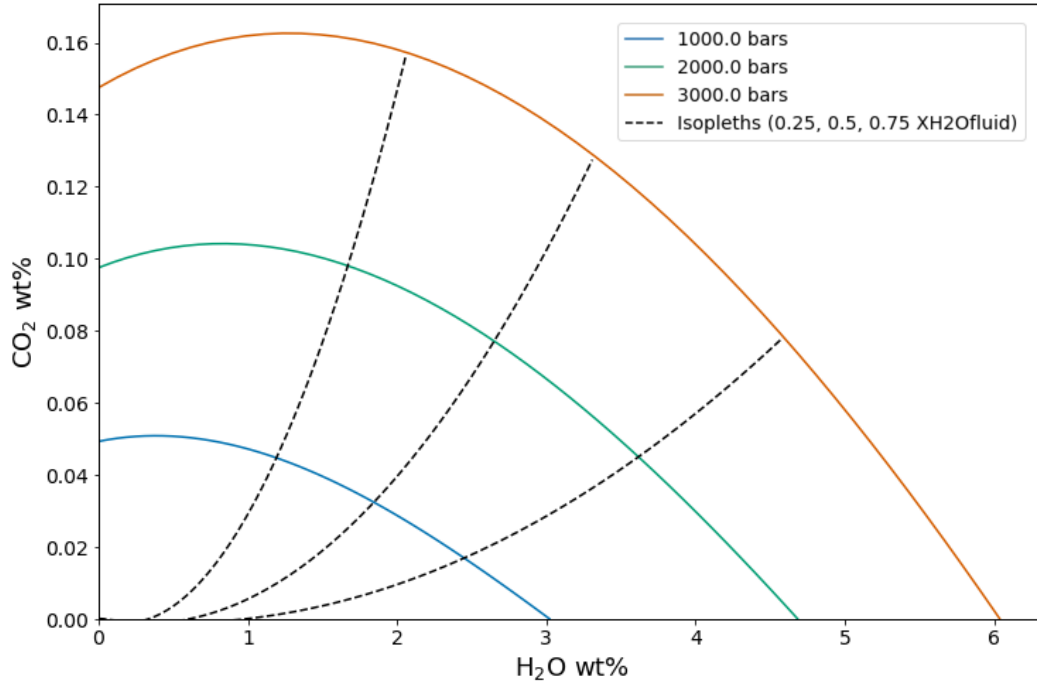
1493 Here we plot the isobars at 1,000, 2,000, and 3,000 bars and isopleths at 0.25,  
 1494 0.5, and 0.75 XH<sub>2</sub>O<sup>fluid</sup> calculated for sample ‘10\*’ at 1,200 °C in Section 3.6 onto  
 1495 one plot.

1496 Input

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1500

```
fig, ax = v.plot(isobars=isobars, isopleths=isopleths)
v.show()
```

1501 Output



**Figure 6.** Isobars and isopleths calculated for the sample, temperature, pressures,  $\text{XH}_2\text{O}^{fluid}$  values, and with the model as defined in Section 3.6. Manuscript default values are sample ‘10\*’ at a 1,200 °C with isobars at 1,000, 2,000, and 3,000 bars, isopleths at  $\text{XH}_2\text{O}^{fluid} = 0, 0.25, 0.5, 0.75,$  and 1 calculated with MagmaSat

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When plotting isobars and isopleths via MagmaSat, the values calculated by `calculate_isobars_and_isopleths()` are used to calculate polynomial fits using numpy’s ‘polyfit’. These polynomial fits, not the raw calculated data, are what have been plotted above. This method of fitting polynomial curves to these data is common in the literature (e.g., Newman & Lowenstern, 2002; IaconoMarziano et al., 2012; Iacovino et al., 2013) and is likely a very close approximation of the true saturation surface. Non-MagmaSat models do not calculate polynomial fits by default, but this can be done by passing `smooth_isobars=True` and `smooth_isopleths=True` to `plot()`.

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A user may wish to apply custom formatting to the plot, in which case the polynomial fits can be calculated and returned as a pandas DataFrame, which the user can then plot up manually using Matplotlib, MS Excel, or some other preferred method. To calculate polynomial fits to isobar and isopleth data, isobars and isopleths can be passed to `smooth_isobars_and_isopleths()`. For this advanced case, we refer the reader to the documentation.

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### 3.8.2 A simple example: Degassing paths

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Here we plot all four degassing paths calculated for sample ‘10\*’ at 1,200 °C in Section 3.7 onto one plot. We designate labels of “Open”, “Half”, “Closed”, and “Exsolved” for the legend.

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Input

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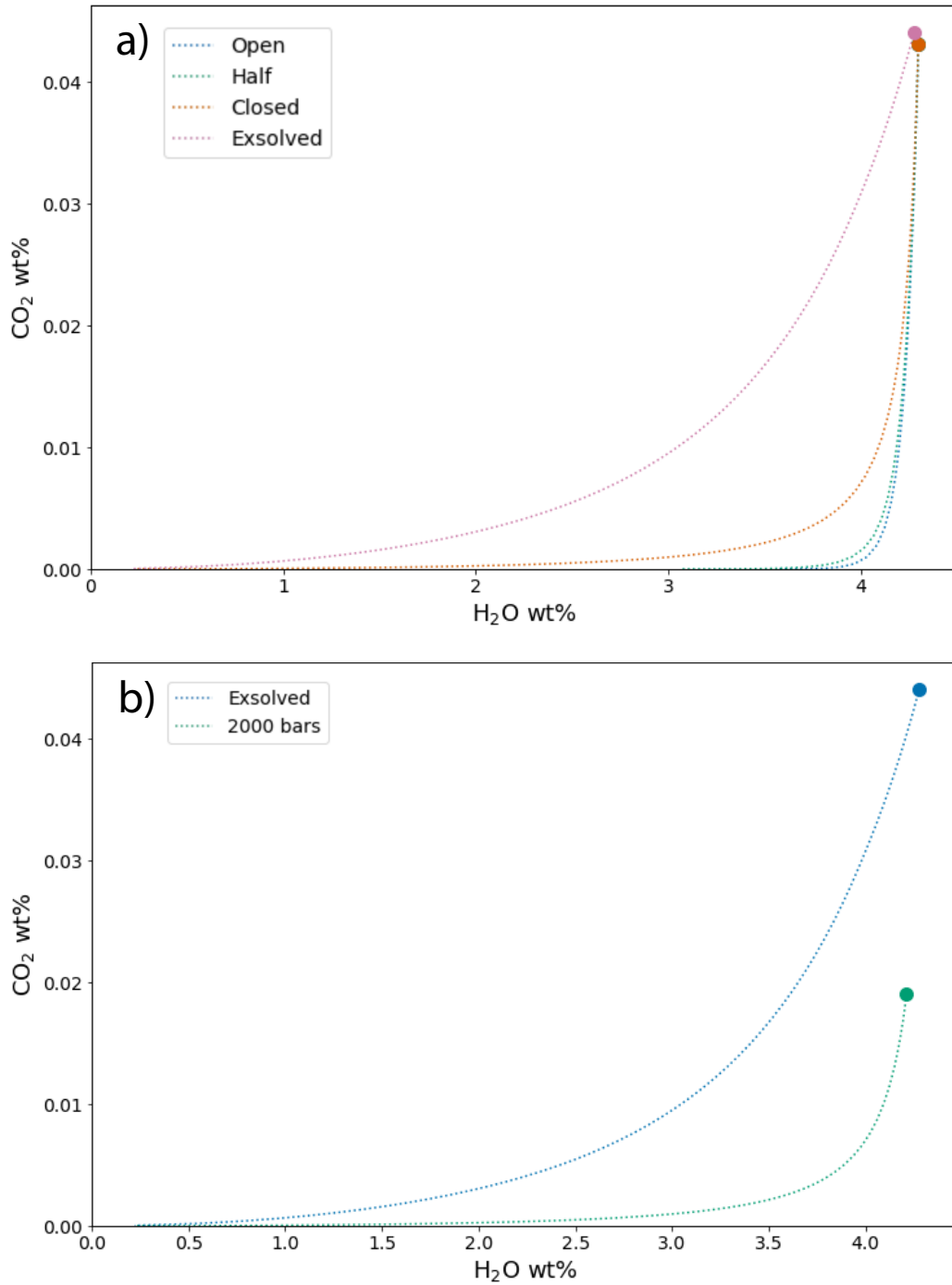
1526

1527

```
fig, ax = v.plot(degassing_paths=[open_df, half_df, closed_df, exsolved_df],
                 degassing_path_labels=["Open", "Half",
                                       "Closed", "Exsolved"])
v.show()
```

1528

Output



**Figure 7.** Degassing paths calculated for the sample, temperature, degassing style, initial exsolved fluid wt%, starting pressure, and model as designated in Section 3.7. Default manuscript values are sample ‘10\*’ at 1,200 °C. “Open”, “Half”, and “Closed” curves in (a) represent open-system, partially open-system (50% fractionated fluid), and closed-system degassing paths, respectively, starting at the saturation pressure. The “Exsolved” curve in (b) represents closed-system degassing with an initial exsolved fluid wt% = 2.0. The “2000” curve in (b) represents closed-system degassing calculated starting at a pressure of 2,000 bars.

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### 3.8.3 Plotting multiple calculations

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One of the major advantages to VESICAL over any other modeling tool is the ability to quickly calculate and plot multiple calculations. VESICAL's `plot()` function is built on top of the popular Matplotlib python library and is designed to work with any VESICAL generated data. It can automatically plot and label one or multiple calculations. In addition, it can plot, as a scatter plot, any x-y points. The `plot()` function always generates plots with H<sub>2</sub>O on the x-axis and CO<sub>2</sub> on the y-axis. `scatterplot()` will take in and plot any x-y data with custom x- and y-axis labels. Generating other commonly used petrologic plots (e.g. Harker style diagrams) is already possible with Matplotlib, and so VESICAL does not duplicate this functionality, however this may be added in future updates.

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It may be tempting to plot multiple calculations on multiple samples and compare them, however we strongly caution against plotting data that do not correspond. For example, isobars and isopleths are calculated isothermally. If degassing paths are also plotted, the user should ensure that the degassing paths were calculated at the same temperature as the isobars and isopleths.

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*3.8.3.1 Isobars, isopleths, and degassing paths* In this example we will use data imported in Section 3.1 and calculations performed in Sections 3.5 and 3.6. Of course, all of the data calculated with VESICAL can be exported to an Excel or CSV file for manipulation and plotting as desired. However, some examples of plotting that can be done within this notebook or in a python script are shown below. Here we plot:

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- Isobars calculated at 1,200 °C and pressures of 1,000, 2,000, and 3,000 bars for sample 10\*
- Isopleths calculated at 1200 °C and XH<sub>2</sub>O<sup>fluid</sup> values of 0, 0.25, 0.5, 0.75, and 1 for sample 10\*
- An open-system degassing path for sample 10\*
- A closed-system degassing path for sample 10\*

1557

Input

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1561

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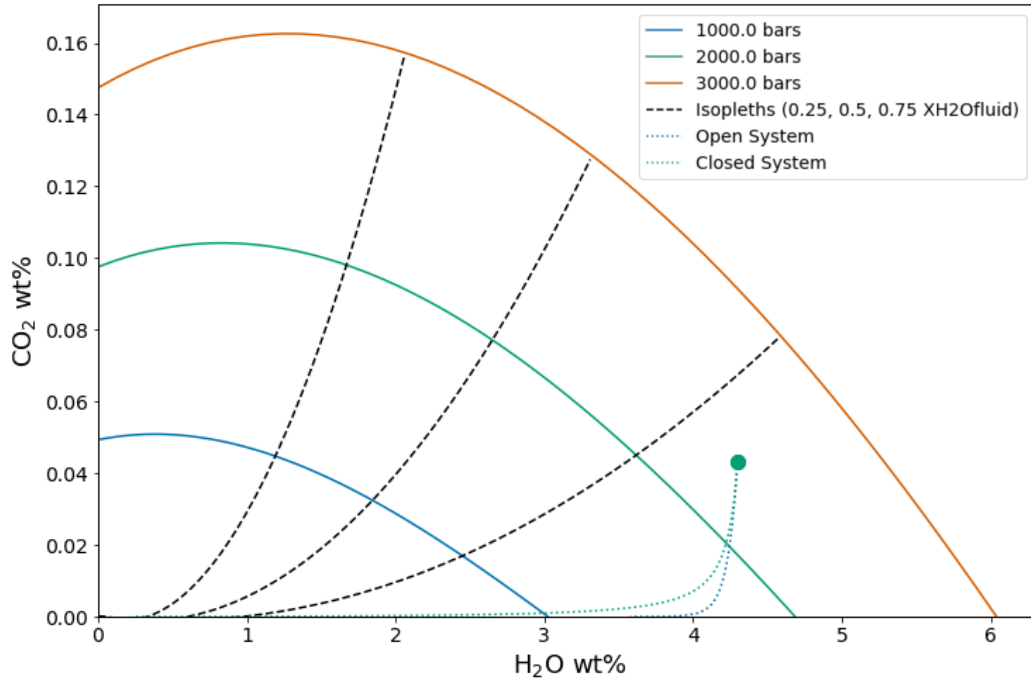
1563

1564

```
fig, ax = v.plot(isobars=isobars, isopleths=isopleths, degassing_paths=[
    open_df, closed_df],
    degassing_path_labels=["Open System",
    "Closed System"])
v.show()
```

1565

Output



**Figure 8.** Example of plotting multiple calculations on one plot. Isobars and isopleths as defined in Section 3.6 and shown in Section 3.8.1 and degassing curves as defined in Section 3.7 and shown in Section 3.8.2. Default manuscript values are for sample ‘10\*’ \*at 1,200 °C with isobars at 1,000, 2,000, and 3,000 bars, isopleths at  $X_{H_2O}^{fluid}$  values of 0, 0.25, 0.5, 0.75, and 1 with an open-system and a closed-system degassing path.

1566 *3.8.3.2 Isobars, isopleths, and degassing paths for multiple samples* First,  
 1567 we will calculate some new data for two different samples: a basanite (sample KI-07  
 1568 from Iacovino et al., 2016) and a rhyolite (sample samp. P1968a from Myers et al.,  
 1569 2019). For both samples we will calculate and then plot:

- 1570 • Isobars and isopleths at 1100 °C, pressures of 1,000 and 2,000 bars and fluid
- 1571 compositions of  $X_{H_2O}^{fluid}$  of 0.25, 0.5, and 0.75
- 1572 • Closed-system degassing paths at 1100 °C

1573 Input:

```

1574 basanite_sample = myfile.get_sample_composition('KI-07', asSampleClass=True)
1575 rhyolite_sample = myfile.get_sample_composition('samp. P1968a', asSampleClass
1576 =True)
1577
1578
1579 basanite_isobars, basanite_isopleths = v.calculate_isobars_and_isopleths(
1580     sample=basanite_sample, temperature=
1581     1100, pressure_list=[1000, 2000],
1582     isopleth_list=[0.25,0.75]).result
1583
1584 rhyolite_isobars, rhyolite_isopleths = v.calculate_isobars_and_isopleths(
1585     sample=rhyolite_sample, temperature=
1586     1100, pressure_list=[1000, 2000],
1587     isopleth_list=[0.25,0.75]).result
  
```

```

1588
1589 basanite_degassing_path = v.calculate_degassing_path(sample=basanite_sample,
1590                                                       temperature=1100).result
1591
1592 rhyolite_degassing_path = v.calculate_degassing_path(sample=rhyolite_sample,
1593                                                       temperature=1100).result
1594

```

1595 Output:

```

1596 Calculating isobar at 1000 bars
1597 done.
1598 Calculating isobar at 2000 bars
1599 done.
1600 Done!
1601 Calculating isobar at 1000 bars
1602 done.
1603 Calculating isobar at 2000 bars
1604 done.
1605 Done!
1606 [=====] 100% Calculating degassing path...
1607 [=====] 100% Calculating degassing path...
1608
1609

```

1610 Input:

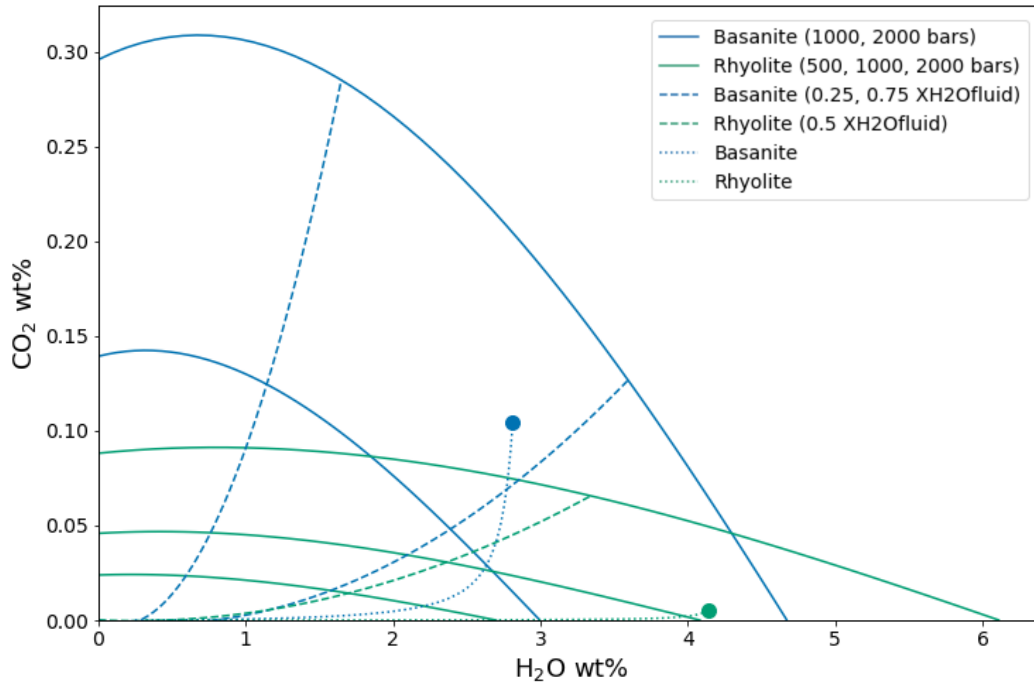
```

1611
1612 fig, ax = v.plot(isobars=[basanite_isobars, rhyolite_isobars],
1613                isopleths=[basanite_isopleths, rhyolite_isopleths],
1614                degassing_paths=[basanite_degassing_path,
1615                                 rhyolite_degassing_path],
1616                isobar_labels=["Basanite", "Rhyolite"],
1617                isopleth_labels=["Basanite", "Rhyolite"],
1618                degassing_path_labels=["Basanite", "Rhyolite"])
1619 v.show()
1620

```

1621 Output:





**Figure 9.** Example of plotting multiple calculations from multiple samples on the same plot. Note that the colors are automatically set to correspond to each sample for all plotted items (here, isobars, isopleths, and degassing paths). Samples, pressures, temperatures,  $X_{H_2O}^{fluid}$  values, and degassing path styles are defined above in this section. Manuscript default values are for a basanite (sample KI-07) and a rhyolite (sample samp. P1968a) at 1,100 °C, 1,000 and 2,000 bars, and  $X_{H_2O}^{fluid} = 0.25$  and 0.75 and closed-system degassing.

1622

### 3.9 Model hybridization (Advanced)

1623

1624

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1630

One of the advantages of implementing the solubility models in a generic python module is the flexibility this affords the user in changing the way solubility models are defined and used. In particular, the structure allows any combination of pure fluid models to be used together in modeling mixed fluids, and fugacity or activity models can be quickly changed without modifying code. This allows advanced users to see how changing a fugacity or activity model implemented in any particular solubility model would affect model results. Instructions for hybridizing models can be found in Supplemental Jupyter notebook S10.

1631

### 3.10 Exporting data

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Once batch calculations have been performed, they can be exported to an Excel or CSV file with the `save_excel()` and `save_csv()` commands. These operations require that the user define a filename (what to name your new file) and a list of the calculation results to save to this file or files.

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Note that this requires that calculations have been assigned to variable names, which has been done in all of the given examples. For example, to calculate saturation pressures of an imported file saved to the variable 'myfile' and simply print the output, the user can type `myfile.calculate_saturation_pressures([options])`, where `[options]` are the required and optional inputs. However, to save this result to a variable (e.g., called 'my\_satPs') so that it can be accessed later, the correct

python syntax would be `my_satPs = myfile.calculate_saturation_pressures([options])`.

Multiple calculations can be saved at once. If saving to an Excel file, each calculation is saved as its own sheet within a single file. If desired, the user can define the names of each of these sheets. If not specified, the sheets will be named 'Original\_User\_Data', which contains the original input data, and then 'CalcN' where N is the nth calculation in a list of calculations. If saving multiple calculations to a CSV file, each calculation will be saved to its own CSV file, and a file name for each of these is required.

Advanced users note that the `calculations` argument takes in any pandas DataFrame object, meaning this functionality is not limited to VESICAL's prescribed outputs. The `save_excel()` and `save_csv()` methods use the pandas `to_excel` and `to_csv` methods, however not all options are implemented here. If saving to a CSV file, any arguments that can be passed to pandas `to_csv` method may be passed to VESICAL's `save_csv()`.

#### Method structure:

```
save_excel(filename, calculations, sheet_name=None)
```

```
save_csv(filenamees, calculations)
```

#### `save_excel()` Required inputs:

**filename** (Excel): Name of the file to create. The extension (.xlsx) should be included along with the name itself, all in quotes (e.g., `filename='myfile.xlsx'`).

**calculations**: A list of variables containing calculated outputs from any of the core BatchFile functions: `calculate_dissolved_volatiles()`, `calculate_equilibrium_fluid_comp()`, and `calculate_saturation_pressure()`. This must be passed as a list type variable, even if only one calculation is given. This is done by enclosing the variable in square brackets (e.g., `calculations=[my_calculation]`).

#### `save_excel()` Optional inputs:

**sheet\_name**: The default value is None, in which case sheets will be saved as 'Original\_User\_data' (the data input by the user) followed by 'CalcN' where N is the nth calculation in `calculations`. Otherwise, a list of names for the sheets can be passed, with the names in quotes (e.g. `sheet_name=['SaturationPressures']`). 'Original\_User\_data' will always be saved as the first sheet.

#### `save_csv()` Required inputs:

**filenamees** (CSV): Name of the file or files to create. The extension (.csv) should be included. If more than one filename is passed, it should be passed as a list. This is done by enclosing the filenames in square brackets (e.g., `filenamees=['`file1.csv'`, ``file2.csv']`).

1682            calculations: same as for `save_excel()`. Must be same length as `filenames`  
 1683            .

### 1684 **Calculated outputs:**

1685            An Excel or CSV file or files will be saved to the active directory (i.e., the  
 1686            same folder as this manuscript notebook or wherever the code is being used).

1687            Here we save five of the calculations performed on an imported data file earlier  
 1688            in this manuscript. The original user-input data are stored in the BatchFile object  
 1689            ‘myfile’. In the following line we use the method `save_excel()` to save the original  
 1690            data and a list of calculations given by the `calculations` argument to an Excel file.

#### 1691 **Input**

```
1692 myfile.save_excel(filename='testsave.xlsx',
1693                  calculations=[dissolved, eqfluid, eqfluid_wtemps, satPs
1694                               , satPs_wtemps],
1695                  sheet_name=['dissolved', 'eqfluid', 'eqfluid_wtemps', '
1696                               SaturationPs', '
1697                               SatPs_wtemps'])
1698
1699
```

#### 1700 **Output**

```
1701 Saved testsave.xlsx
1702
1703
```

### 1704 **3.10.1 Saving data for re-import into VESICAL**

1705            In many cases, it may be preferable to compute large amounts of data using  
 1706            VESICAL and then reimport them, either to perform more analysis or to plot the  
 1707            data. Likewise, a user may wish to compute data in VESICAL and then send the re-  
 1708            sults to a colleague, who can then re-import that data into VESICAL directly. For  
 1709            this case, we suggest using python’s pickle package ([https://wiki.python.org/  
 1710            moin/UsingPickle](https://wiki.python.org/moin/UsingPickle)). Any python object, such as the results of a VESICAL calcula-  
 1711            tion, can be “pickled” or saved as a python-readable file. To use pickle, users must  
 1712            first import the pickle module, then “dump” the desired contents to a pickle file.  
 1713            The pickled data can be accessed by “loading” the pickled file.

1714            Below we pickle our computed dissolved volatile concentrations by dumping  
 1715            our variable `dissolved` to a pickle file that we name “dissolved.p”.

```
1716 import pickle
1717
1718
1719 pickle.dump(dissolved, open("dissolved.p", "wb"))
1720
```

1721            In another python file or terminal session, `dissolved` can be loaded back in via:

```
1722 import pickle
1723
1724
1725 dissolved = pickle.load(open("dissolved.p", "rb"))
1726
```

## 1727 4 Discussion and Applications

### 1728 4.1 Compositional Variation Within Datasets and Best Practices

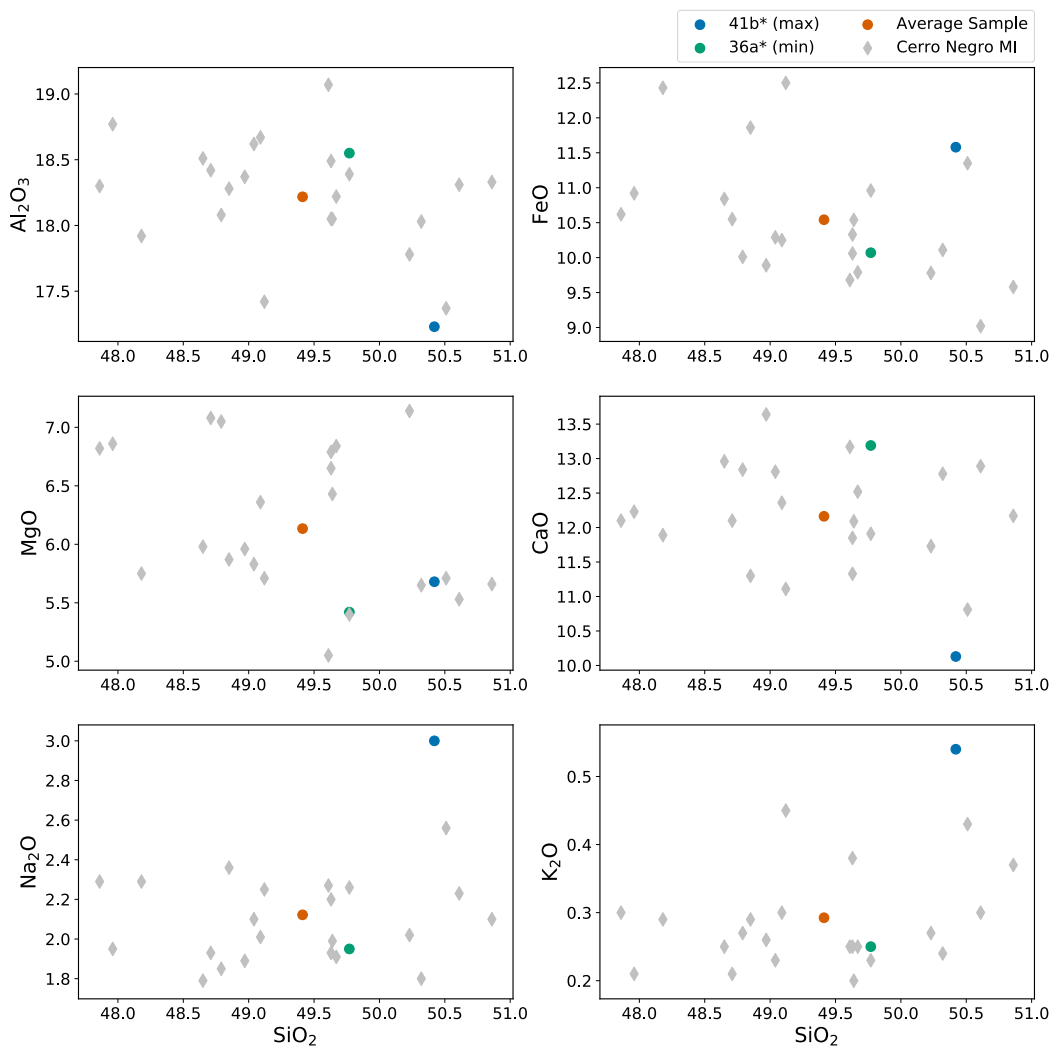
1729 While not all solubility models incorporate significant bulk compositional pa-  
 1730 rameters, it has been clearly shown that the composition of a melt plays a strong  
 1731 role in determining the solubility of H<sub>2</sub>O and CO<sub>2</sub> in magmas (Ghiorso & Gualda,  
 1732 2015; Moore, 2008; Papale et al., 2006; Wieser et al., submitted). Thus, composi-  
 1733 tional variance must be accounted for in any study examining solubility in multiple  
 1734 samples. A key use case where VESIcal can facilitate the adoption of this practice  
 1735 is in melt inclusion (MI) studies; specifically, where a single suite of MI with multi-  
 1736 ple melt compositions is examined using solubility models to interrogate magmatic  
 1737 degassing processes. Prior to the availability of VESIcal, the difficulty associated  
 1738 with performing multiple model calculations on multiple samples resulted in very  
 1739 few studies accounting for any compositional variance within their datasets. Indeed,  
 1740 until now, it has been difficult to even assess whether the potentially minimal com-  
 1741 positional variance within a suite of melt inclusions from a single volcanic eruption  
 1742 would have any measurable effect on solubilities calculated for different MI.

1743 Using VESIcal, we can address the question: what is the quantitative effect of  
 1744 compositional variation within a single suite of melt inclusions upon calculated melt  
 1745 inclusion saturation pressures? And, how does this affect conclusions that might  
 1746 be drawn regarding volcanic degassing and eruptive processes? To investigate this,  
 1747 we use a dataset of basaltic melt inclusions from Cerro Negro volcano, Nicaragua  
 1748 (Roggensack, 2001). The compositional variation of these MI (Figure 10), while rela-  
 1749 tively restricted, results in quite variable mixed-fluid solubilities from sample to sam-  
 1750 ple. To determine the end-member compositions within the dataset corresponding to  
 1751 the samples with the maximum and minimum combined H<sub>2</sub>O-CO<sub>2</sub> solubilities, iso-  
 1752 bars were computed at 1200 °C and 3,000 bars for all samples using the MagmaSat  
 1753 model in VESIcal. Maximum and minimum samples were taken as the isobar curves  
 1754 with the smallest and largest integral (area under the curve). We refer to this value  
 1755 as the “integrated mixed-volatile solubility” value, IMS, in units of concentration  
 1756 squared. The samples that produced maximum and minimum integrated solubili-  
 1757 ties are shown in Figures 10 and 11 in blue and green, respectively (sample 41b\*,  
 1758 IMS=0.81 and 36a\*, IMS=0.66 wt%<sup>2</sup> at 3,000 bars). A composition representing the  
 1759 average of all MI in the dataset is shown in orange (“Average Sample”, IMS=0.70  
 1760 wt%<sup>2</sup> at 3,000 bars). A jupyter notebook to reproduce these calculations is provided  
 1761 in the supplement (Supplementary Jupyter Notebook S8).

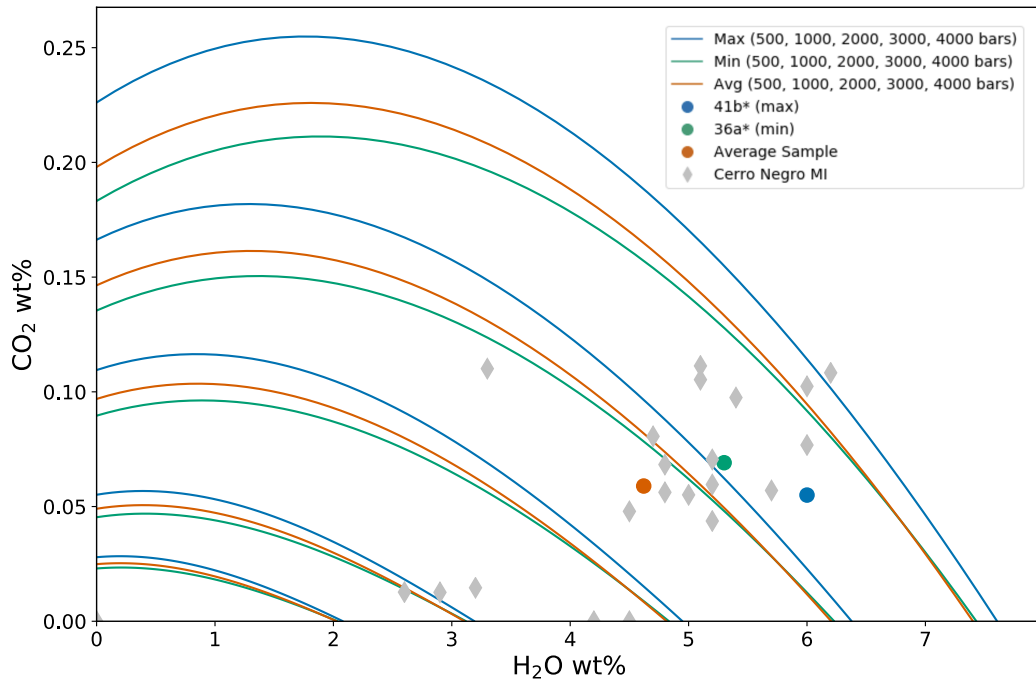
1762 At all pressures, the integrated mixed-volatile solubility across the Cerro Ne-  
 1763 gro dataset varies as much as 10% relative (Figure 11). For these MI, this results in  
 1764 as much as 11.5% relative error in the calculation of saturation pressures (average  
 1765 error for the entire dataset of 6.8% relative). It is noteworthy that this error is not  
 1766 systematic either in terms of absolute value or sign. For example, when calculated  
 1767 using their own compositions, saturation pressures for maximum and minimum sam-  
 1768 ples 41b\* and 36a\* are 3050 and 3090 bars, respectively. But, saturation pressures  
 1769 calculated for both of these MI using the dataset’s average composition are 3020  
 1770 and 3250 bars, respectively. That is an error of -30 and +160 bars or -1% and +5%  
 1771 respectively. Errors in these calculations, thus, may be quite small. But, in any case,  
 1772 removing this error completely is a simple task using VESIcal, and so we recommend  
 1773 that studies adopt the practice of calculating volatile solubilities (and associated  
 1774 values) in melts using the composition unique to each melt investigated.

1775 Even in cases where solubility values (e.g., saturation pressures) are not calcu-  
 1776 lated, the error highlighted above plagues any isobar diagram over which multiple  
 1777 melt compositions are plotted (e.g., Figure 11). Alternative plots to the commonly  
 1778 used H<sub>2</sub>O-CO<sub>2</sub> diagram are shown in Figure 12, in which the same dataset is plot-

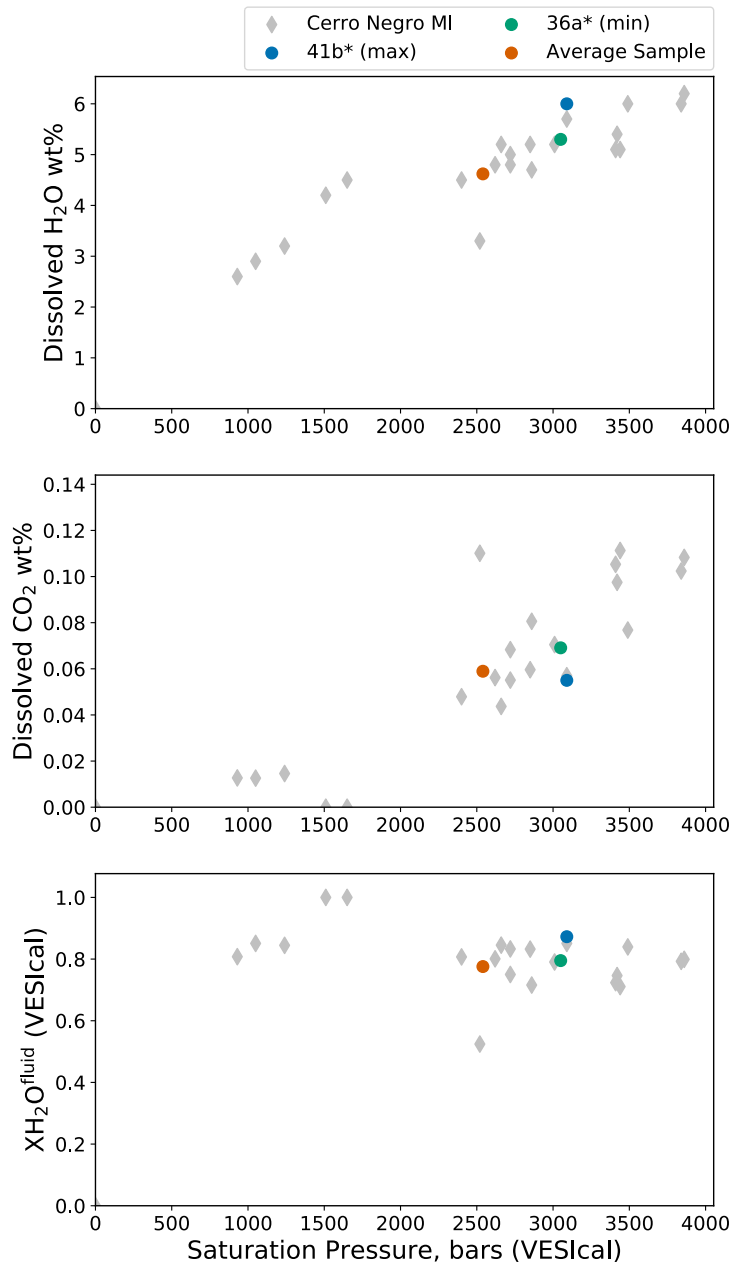
1779 ted in terms of computed saturation pressure (at 1200 °C calculated with VESICAL  
 1780 using MagmaSat) versus dissolved H<sub>2</sub>O, dissolved CO<sub>2</sub>, and fluid composition (as  
 1781 XH<sub>2</sub>O<sup>fluid</sup> calculated with VESICAL using MagmaSat). These plots avoid the is-  
 1782 sues discussed above as they are compositionally independent, since the saturation  
 1783 pressure is calculated individually for each sample composition. Degassing trends  
 1784 are more accurately represented; H<sub>2</sub>O and CO<sub>2</sub> concentrations lie along expected  
 1785 degassing trends with much less scatter than the H<sub>2</sub>O-CO<sub>2</sub> plot. We can also see  
 1786 from this figure that the fluid composition during this eruption at Cerro Negro re-  
 1787 mained relatively constant at XH<sub>2</sub>O<sup>fluid</sup> ~0.8 from reservoir to surface, suggesting  
 1788 a scenario approaching closed-system degassing (i.e., melt volatile concentrations are  
 1789 buffered by the co-existing fluid composition).



**Figure 10.** Harker style diagrams illustrating the compositional range of MIs from Cerro Negro volcano from Roggensack (2001). The “Average Sample” plotted as an orange dot represents a fictitious sample, calculated as the average of all MIs in the dataset. Sample 41b\* and 36a\* are the names of samples that produced isobars with maximum and minimum area under the curve, respectively (see text). Gray diamonds are all other data in the dataset.



**Figure 11.** H<sub>2</sub>O-CO<sub>2</sub> diagram with isobars for MI from Cerro Negro volcano (Roggensack, 2001) computed by VESICAL using MagmaSat at 1200 °C, pressures of 500, 1000, 2000, 3000, and 4,000 bars. Curves shown are polynomials fitted to data computed by VESICAL. Blue and green curves correspond to samples 41b\* and 36a\*, which produced isobars with maximum and minimum area under the curve, respectively. Orange isobars were those computed for a fictitious sample representing the average composition of the MI dataset. Gray diamonds are all other data in the dataset.



**Figure 12.** Saturation pressure at 1200 °C calculated using VESIcal with MagmaSat versus measured dissolved H<sub>2</sub>O and CO<sub>2</sub> concentrations and calculated fluid composition in Cerro Negro melt inclusions. These plots meaningfully illustrate degassing processes while avoiding issues associated with commonly used H<sub>2</sub>O-CO<sub>2</sub> diagrams, which occur with even minor compositional variation within a given dataset.

## 4.2 Model Comparisons

One of the possible workflows enabled through VESICAL is the ability to compute and compare (numerically and graphically) results from several models at once. To illustrate this point, we will take two single samples within the calibrated compositional range of several models, calculate isobars at multiple pressures, and plot the results. This is a common way to compare the solubility surface computed by different models for a single melt composition, and it is particularly useful since it quickly highlights the significant variation that exists between published models. The results of this exercise are shown here, and a Jupyter notebook to reproduce the code and calibration checks is available in the Supplement (Supplementary Jupyter Notebook S9).

We use a fictitious alkali basalt that we name “alkbasalt” and a fictitious rhyolite whose compositions are given in Table 12. The use of VESICAL’s `calib_plot()` function (see supplement) illustrates that the composition of the alkali basalt is within the compositional calibration ranges of four mixed-fluid solubility models: MagmaSat, Iacono-Marziano, Dixon, and ShishkinaIdealMixing. The rhyolite is within the ranges of MagmaSat and Liu. Isobars were calculated with these models at 1200 °C for alkbasalt and 800 °C for rhyolite and pressures of 500, 1,000, and 2,000 bars, using the below code:

Input

```
model_comps = v.BatchFile("Table_Model_Comps.xlsx")
model_comps.get_data()
```

**Table 12.** Melt compositions used for modeling

	SiO2	TiO2	Al2O3	Fe2O3	FeO	MnO	MgO	CaO	Na2O	K2O	P2O5	H2O	CO2	Cr2O3	NiO	CoO
<b>Alkali Basalt</b>	49.00	1.27	19.7	3.74	5.33	0.17	4.82	8.85	4.23	1.00	0.37	4.51	0.25	0.0	0.0	0.0
<b>Rhyolite</b>	77.19	0.06	12.8	0.00	0.94	0.00	0.03	0.53	3.98	4.65	0.00	0.26	0.05	0.0	0.0	0.0

Input

```
alkbasalt = model_comps.get_sample_composition("Alkali Basalt", asSampleClass=True)
rhyolite = model_comps.get_sample_composition("Rhyolite", asSampleClass=True)

alkbasalt_isobars, alkbasalt_isopleths = v.calculate_isobars_and_isopleths(
    sample=alkbasalt, temperature=1200,
    pressure_list=[500, 1000, 2000],
    isopleth_list=[0.5], print_status=True
).result

rhyolite_isobars, rhyolite_isopleths = v.calculate_isobars_and_isopleths(
    sample=rhyolite, temperature=800,
    pressure_list=[500, 1000, 2000],
    isopleth_list=[0.5]).result

Iac_alkbasalt_isobars, Iac_alkbasalt_isopleths = v.
    calculate_isobars_and_isopleths(sample
```



```

1833         =alkbasalt, temperature=1200,
1834         pressure_list=[500, 1000, 2000],
1835         isopleth_list=[0.5], model="
1836         IaconoMarziano").result
1837
1838 Dixon_alkbasalt_isobars, Dixon_alkbasalt_isopleths = v.
1839         calculate_isobars_and_isopleths(sample
1840         =alkbasalt, temperature=1200,
1841         pressure_list=[500, 1000, 2000],
1842         isopleth_list=[0.5], model="Dixon").
1843         result
1844
1845 Shish_alkbasalt_isobars, Shish_alkbasalt_isopleths = v.
1846         calculate_isobars_and_isopleths(sample
1847         =alkbasalt, temperature=1200,
1848         pressure_list=[500, 1000, 2000],
1849         isopleth_list=[0.5], model="
1850         ShishkinaIdealMixing").result
1851
1852 Liu_rhyolite_isobars, Liu_rhyolite_isopleths = v.
1853         calculate_isobars_and_isopleths(sample
1854         =rhyolite,
1855         temperature=800, pressure_list=[500, 1000, 2000], isopleth_list=[0.5], model=
1856         "Liu").result
1857

```

## Output

```

1858 Calculating isobar at 500 bars
1859 done.
1860 Calculating isobar at 1000 bars
1861 done.
1862 Calculating isobar at 2000 bars
1863 done.
1864 Done!
1865 Calculating isobar at 500 bars
1866 done.
1867 Calculating isobar at 1000 bars
1868 done.
1869 Calculating isobar at 2000 bars
1870 done.
1871 Done!
1872
1873 RuntimeWarning: pressure exceeds 1000 bar, which Iacono-Marziano et al. (2012)
1874 suggest as an upper calibration limit of the Dixon (1997, Pi-SiO2 simpl.)
1875 Model
1876
1877
1878

```

## Input

```

1879 fig, ax = v.plot(isobars=[alkbasalt_isobars, Iac_alkbasalt_isobars,
1880 Dixon_alkbasalt_isobars,
1881 Shish_alkbasalt_isobars],
1882 isobar_labels=["MagmaSat", "Iacono-
1883 Marziano", "Dixon", "Shishkina"])
1884
1885 v.show()
1886
1887

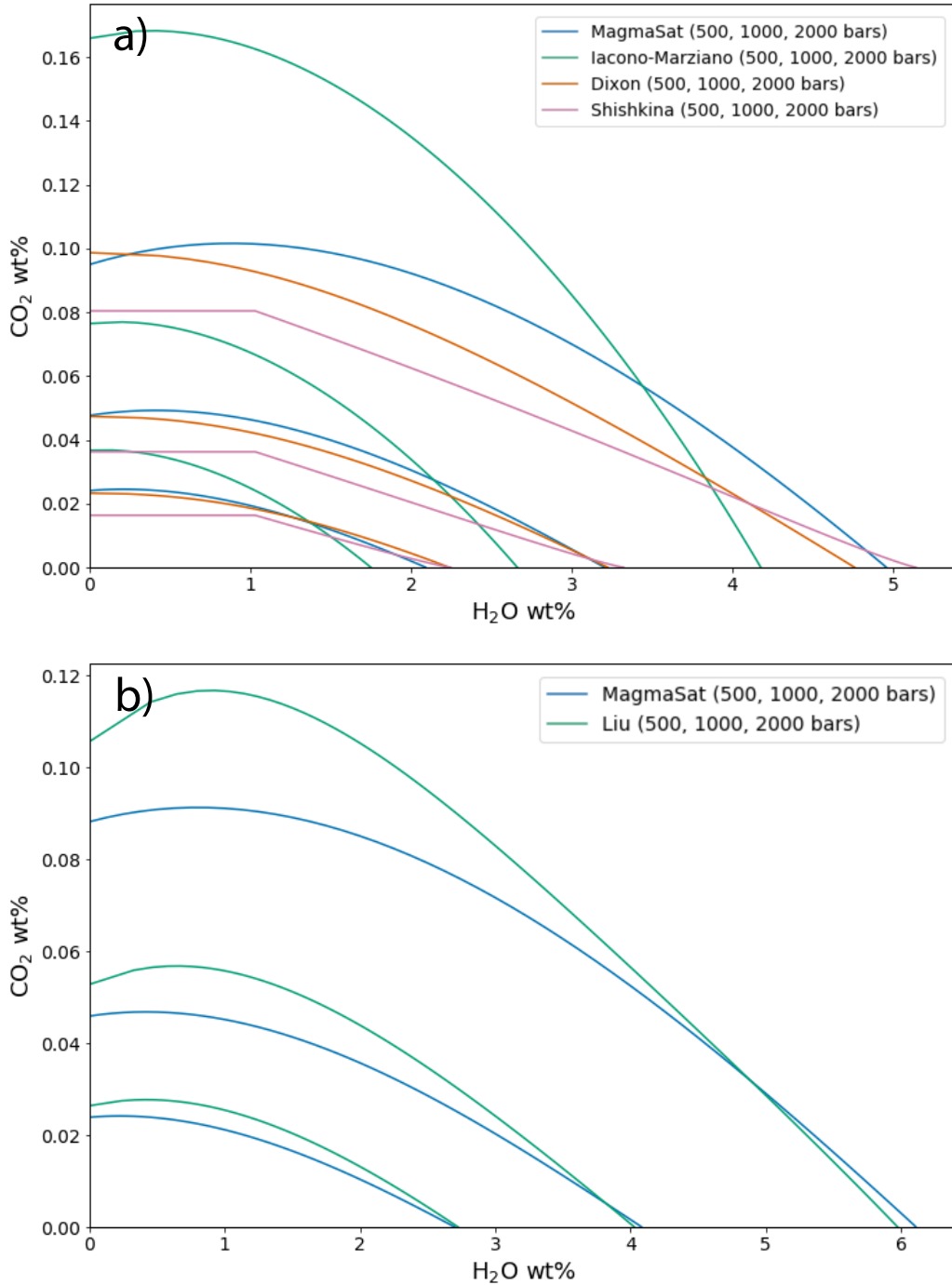
```

## Output

```

1888

```



**Figure 13.** Isobars plotted for an alkali basalt (a) and rhyolite (b) with VESICAL for five mixed-fluid solubility models. For alkali basalt, MagmaSat, Iacono-Marziano, Dixon, and ShishkinaIdealMixing were used to create isobars at 1,200 °C. For rhyolite, MagmaSat and Liu were used to create isobars at 800 °C.

1889  
1890  
1891

It is immediately clear from Fig. 13 that major disagreement exists between these models. For the alkali basalt, MagmaSat and Dixon show the best agreement, particularly at pressures <2000 bars. However, the mismatch between these models

(and, indeed, between all models) increases with pressure. The Iacono-Marziano model is calibrated for highly depolymerized alkali basalts resulting in an increased capacity of the melt to dissolve  $\text{CO}_3^{2-}$ . That may explain why this model predicts significantly higher  $\text{CO}_2$  solubilities at  $\text{XH}_2\text{O}^{\text{fluid}}$  values approaching 0.

The ShishkinaIdealMixing model displays nearly linear isobars, with finite solubility below  $\sim 1$  wt% dissolved  $\text{H}_2\text{O}$ . This is a consequence of the model calibration; the pure- $\text{H}_2\text{O}$  solubility expression of ShishkinaIdealMixing is not calibrated with any experiments at low  $\text{PH}_2\text{O}$ . This results in a finite solubility at low dissolved  $\text{H}_2\text{O}$  concentrations, such that the zero-pressure solubility is not zero. This produces significant model error at low but non-zero values of  $\text{XH}_2\text{O}^{\text{fluid}}$ . Thus, we caution the user against using the Shishkina model at low but non-zero  $\text{XH}_2\text{O}^{\text{fluid}}$  or when fluids deviate far from pure  $\text{H}_2\text{O}$  or pure  $\text{CO}_2$ . In general, the Shishkina model should only be used for modeling pure- $\text{H}_2\text{O}$  or pure- $\text{CO}_2$  fluids. This is discussed in more detail in Wieser et al. (submitted).

The models of MagmaSat and Liu show a similar level of disagreement for  $\text{H}_2\text{O}$ - $\text{CO}_2$  solubility in the rhyolite, with Liu predicting much higher dissolved  $\text{CO}_2$  concentrations at low  $\text{XH}_2\text{O}^{\text{fluid}}$  ( $< 20$  relative% or  $\sim 220$  ppm at  $\text{XH}_2\text{O}^{\text{fluid}}=0.1$ ).

### 4.3 Sensitivity and error analysis

To date, very few studies have compared the sensitivity of their pressure estimates to the choice of solubility model, or propagated errors inherent to measurements of volatile concentrations in melts using SIMS, FTIR and Raman Spectroscopy into an error bar in terms of saturation pressure. In contrast, VESICAL allows users to import an Excel or CSV spreadsheet with each row containing the major element and volatile contents of each inclusion, as well as a temperature at which to evaluate solubility. Using the batch calculation functions, VESICAL will automatically calculate the saturation pressure for each row, using a user-specified model. Thus, users can more easily compare results from different solubility models, to robustly assess their applicability for the system of interest. Additionally, users could load a different spreadsheet, where the  $\text{CO}_2$  and  $\text{H}_2\text{O}$  concentrations are adjusted to reflect the analytical uncertainty on the instrument used, allowing error bars on the saturation pressure to be calculated for every single inclusion. The open-source nature of VESICAL also allows the capabilities of Python3 to be applied to modeling of solubility in magmas. For example, users could utilize Markov chain Monte Carlo (MCMC) methods to robustly calculate error distributions for each sample. In future releases, automatic sensitivity and error analysis on datasets and calculated results may be implemented directly within VESICAL, building on existing tools within the python community.

### 4.4 Future development

VESICAL represents the first comprehensive volatile solubility modeling tool of its kind, including the feature that VESICAL is extensible. VESICAL is written so that implementing new or yet-to-be-implemented solubility models is as simple as possible. To implement a new model, python code describing the model equations needs to be written, and this model name needs to be added to a list of model names within the code. To make this as simple as possible such that the original authors of VESICAL are not the only people who can develop the code, planned future work includes the creation of detailed instructions (including instructional videos) illustrating this process.

Likewise, new features can be added at any time, and enthusiastic members of the community who wish to help bring such features to VESICAL are very welcome.

1941 Users can contribute to VESIcal’s code, implementing new models and new fea-  
 1942 tures, via github (<https://github.com/kaylai/VESIcal>). The repository is public,  
 1943 but we encourage users who wish to contribute to the code to fork the repository  
 1944 into their private workspace on github. Once edits to the code are complete, the  
 1945 new code can be added to VESIcal by creating a “Pull Request” inside of github.  
 1946 Changes and enhancements to VESIcal will correspond to a change in the code’s  
 1947 version number. The published version of the code documented in this manuscript  
 1948 and archived on Zenodo is version 0.9.10 (DOI: 10.5281/zenodo.4652839). Planned  
 1949 features not implemented in this release include: 1. Models to calculate sample oxy-  
 1950 gen fugacity from  $\text{Fe}^{2+}/\sum\text{Fe}$  and vice versa; 2. Additional volatiles such as sulfur;  
 1951 3. More thermodynamic solubility models such as that of Papale et al. (2006); 4.  
 1952 Sensitivity and error analysis functions.

#### 1953 4.5 How to cite VESIcal and its models

1954 To cite computations done using VESIcal, please cite this manuscript, the  
 1955 VESIcal version number, as well as the model(s) used. Note that if a model was not  
 1956 specified during calculations, the default model of MagmaSat was used and should  
 1957 be cited as “MagmaSat Ghiorso and Gualda (2015)”. For example: “Calculations  
 1958 were performed using VESIcal (v. 0.9.10; Iacovino et al., 2020) with the models of  
 1959 Shishkina et al. (2014) and Dixon (1997, “VolatileCalc”).” The web-app always runs  
 1960 on the most up-to-date version of the VESIcal code, but it is best practice to note if  
 1961 the web-app was used (“Calculations were performed using the VESIcal web-app  
 1962 (v. 0.9.10; Iacovino et al., 2020)...”). We also encourage users to be as explicit as  
 1963 possible as to the conditions used for modelling. This includes stating the pressure,  
 1964 temperature, volatile concentration, and bulk magma composition used in modelling.  
 1965 In the best case, VESIcal users will provide their code (e.g., as a jupyter notebook  
 1966 or .py file) along with their publication such that it can be easily replicated.

## 1967 5 Conclusions

1968 VESIcal is a thermodynamic mixed-volatile solubility engine designed to meet  
 1969 the growing computational needs of the igneous petrology community. Seven com-  
 1970 monly used volatile solubility models are built into VESIcal, which employs the most  
 1971 diversely calibrated (chemically and in P-T space) of the group, MagmaSat (Ghiorso  
 1972 & Gualda, 2015), as the default model. VESIcal can perform five core calculations  
 1973 with any mixed-fluid model and three core calculations with any model (mixed-fluid,  
 1974  $\text{CO}_2$ -only,  $\text{H}_2\text{O}$ -only). VESIcal allows for automatic calculation of large datasets and  
 1975 robust built-in plotting capability.

1976 Alongside model frameworks such as ENKI, VESIcal represents an early step  
 1977 forward toward creating a generalized thermodynamic framework to model whole  
 1978 scale magmatic processes. Such a framework builds upon the key tenets of VESIcal;  
 1979 namely: fundamental thermodynamic underpinning; inclusion of existing modeling  
 1980 strategies; python powered, open-source, and extensible code base; high usability at  
 1981 all levels; benchmarking and testing; and power as a responsive and predictive tool.

## 1982 Data Availability Statement

1983 The VESIcal software is open source and is hosted on github ([https://](https://github.com/kaylai/VESIcal)  
 1984 [github.com/kaylai/VESIcal](https://github.com/kaylai/VESIcal)). The version of VESIcal used in this manuscript  
 1985 is version 0.9.10 and is archived on zenodo (DOI: 10.5281/zenodo.4652839). VESIcal  
 1986 runs on top of thermoengine, a python package that is a part of the ENKI frame-  
 1987 work (<http://enki-portal.org/>). The thermoengine library is open source and is  
 1988 available on GitLab (<https://gitlab.com/ENKI-portal/ThermoEngine>). VESI-

cal was written in Python3 and should be stable up to at least Python version 3.7.6. In addition to thermoengine, VESICAL requires the following standard libraries (with versions used for testing indicated in brackets): pandas (1.0.1), numpy (1.18.1), matplotlib (3.1.2), cycler (0.10.0), scipy (1.4.1), and sympy (1.5.1). The VESICAL webapp interface runs through Anvil ([anvil.works](https://anvil.works)), which executes VESICAL code on a cloud server. The code that facilitates the link between the anvil interface and the VESICAL code is available on the VESICAL github. VESICAL can also be used within a jupyter notebook and is hosted on the ENKI jupyter hub (<https://server.enki-portal.org/hub/login>) such that the code can be accessed without installation on the user's local machine.

All data sets used in this manuscript are available on the VESICAL github as well as in the Supplementary Material of this manuscript. The example dataset used for worked examples in Section 3 (example\_data.xlsx file; Supplemental Dataset S1) contains compositional information for basalts (Roggensack, 2001; Tucker et al., 2019), andesites (Moore et al., 1998), rhyolites (Mercer et al., 2015; Myers et al., 2019), and alkaline melts (phototephrite, basaltic-trachyandesite, and basanite from Iacovino et al. 2016). Several additional example datasets from the literature are available in the Supplement (Supplementary Datasets S2-S5; Table 4). These include experimentally produced alkaline magmas from Iacovino et al. (2016, alkaline.xlsx), basaltic melt inclusions from Kilauea (Tucker et al., 2019) and Gakkel Ridge (Bennett et al., 2019, basalts.xlsx), basaltic melt inclusions from Cerro Negro volcano, Nicaragua (Roggensack, 2001, cerro\_negro.xlsx), and rhyolite melt inclusions from the Taupo Volcanic Center, New Zealand (Myers et al., 2019) and a topaz rhyolite from the Rio Grande Rift (Mercer et al., 2015, rhyolites.xlsx). Where available, the calibration datasets for VESICAL models are also provided (Supplementary Datasets S6-S7).

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