1	Capturing the spatial variability of algal bloom development in a shallow temperate lake
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11	Abstract
12	1. Algal blooms can have profound effects on the structure and function of aquatic
13	ecosystems and have the potential to interrupt valuable ecosystem services. Despite the
14	potential ecological and economic consequences of algal blooms, the spatial dynamics of
15	bloom development in spatially complex ecosystems such as shallow lakes remain poorly
16	characterized. Our goal was to evaluate the magnitude and drivers of spatial variability of
17	algal biomass, dissolved oxygen and pH over the course of a season, in a shallow lake in
18	order to better understand the spatial dynamics of algal blooms in these ecosystems.
19	2. We sampled 98 locations in a small eutrophic lake on a 65m grid for several parameters
20	(chlorophyll a, phycocyanin, dissolved oxygen, pH, and temperature), weekly over 122
21	days. This was done to estimate the dynamics of variability and spatial autocorrelation
22	during the course of multiple bloom events. We also compared the spatial measurements
23	to a high frequency sensor deployed at a fixed station and estimated the optimal spatial
24	sampling resolution by performing a rarefaction analysis.

25 3. Spatial heterogeneity of algal pigments was high, particularly during bloom events, and 26 this pattern and the overall severity of the bloom was not well captured with the fixed 27 station monitoring. The pattern of algal pigments and other limnologically important 28 variables (dissolved oxygen and pH) was related to the direction of prevailing winds 24 29 hours prior to sampling and likely enhanced by the shallow northern basin of the lake 30 where the main surface inlet was located. Additionally, a dense bed of floating-leaf 31 macrophytes contributed to local patchiness in all variables. Finally, from the rarefaction 32 analysis we found that minimal information about the mean state of the ecosystem was 33 gained after ~30 locations had been sampled.

4. This study revealed how spatially heterogeneous shallow lakes are over the course of a
single season, and that the magnitude of variability was highest during biologicallyintensive periods such as algal blooms. As such, continued research is needed across a
range of trophic conditions to better understand the structure of horizontal variability in
lakes. Overall, these data demonstrate the need for spatially-explicit monitoring to better
understand the dynamics and drivers of algal blooms in shallow lakes and to better
manage ecosystem services.

41 Introduction

42 Lakes are highly dynamic ecosystems that can undergo rapid physical and chemical 43 changes at an individual location, throughout their water column, and across the entire lake 44 surface at the scale of hours, days, seasons, and years (Laas et al., 2012; Read et al., 2011; Wynne & Stumpf, 2015). Quantifying heterogeneity in aquatic ecosystem structure and function 45 46 not only improves our understanding of lake ecology and the underlying mechanisms that drive 47 spatial and temporal heterogeneity, but also provides insights that improve management of these 48 ecosystems and the services they provide. With the development of sophisticated sensor 49 technology, high frequency measurements of variables such as dissolved oxygen and temperature 50 have helped limnologists grasp the scale of temporal heterogeneity in lakes (Carpenter et al., 51 2020; Chaffin et al., 2020; Cotterill et al., 2019). Detailed temporal monitoring has led to 52 advances in understanding several lake mechanisms such as diel cycles in primary production 53 (Solomon et al., 2013; Staehr et al., 2012), temperature effects on biogeochemical processes 54 (Medeiros et al., 2012), and early warnings of the transition to alternative stable states (Carpenter 55 et al., 2011; Wilkinson et al., 2018). Additionally, high frequency measurements have been used 56 to better understand heterogeneity over depth (vertical spatial heterogeneity) for important 57 processes such as stratification (Boehrer & Schultze, 2008; Read et al., 2011). Despite these 58 advances in understanding temporal and vertical heterogeneity, less is known about the dynamics 59 of horizontal spatial heterogeneity in the surface waters of lakes.

The vast majority of our understanding of lentic ecosystem structure and function comes from single station sampling, with measurements taken through time over the deepest point in the lake (Stanley et al., 2019). This location is usually selected to be representative of conditions in the lake; however, the representativeness of a single location is likely to vary with regards to

64 the variable being measured and with time due to interacting forces such as wind, hydrology, 65 bathymetry, and biology (Chaffin et al., 2020; Schilder et al., 2013; Wu et al., 2010; Zhou., 66 2013). For example, ecosystem metabolism measured at dozens of locations for 10 days in two 67 north temperate lakes varied 1-2 orders of magnitude, with more than three-quarters of the 68 variability attributable to the measurement location within the lake (Van de Bogert et al., 2012). 69 Transect-based studies of reservoirs have revealed gradients in algae pigments, pH, and nutrients 70 with differences varying between 25%-180% within a waterbody (Moreno-Ostos et al., 2009; 71 Rychtecky & Znachor, 2011; Smith, 2018). Recently, satellite-based studies have demonstrated 72 the ability to detect spatial patterns at a high resolution for optical variables in large lakes (Lekki et al., 2019). Despite these advances, relatively few studies have quantified horizontal spatial 73 74 variability over time in lakes (Buttita et al. 2017, Vilas et al. 2017, Loken et al. 2019), hampering 75 our understanding of the magnitude of heterogeneity in variables important for managing water 76 quality and ecosystem services. 77 The development of algal blooms is expected to be a spatially heterogeneous 78 phenomenon (Buelo et al., 2018; Butitta, Carpenter et al., 2017; Serizawa et al., 2008) due to 79 both local heterogeneity in nutrient limitation, zooplankton grazing, and temperature (Davis et 80 al., 2009; Hansen et al., 1997) and population scale heterogeneity due to wind (George &

81 Heaney, 1978). Algal blooms can have a negative effect on ecosystem services, and therefore are

82 often a target for ecosystem monitoring and management. Some bloom-forming taxa,

83 particularly freshwater cyanobacteria, can produce toxins that rise to dangerous concentrations

for humans, pets, and livestock (Codd et al., 2005; Corbel et al., 2014). Additionally, the

85 mineralization of settling phytoplankton contributes to anoxic bottom waters, while intense

86 periods of primary production cause large variation in dissolved oxygen and pH (in poorly

buffered ecosystems) over the course of the day, which is stressful for aquatic organisms
(Gilbert, 2017; Landsberg, 2002). Furthermore, the perceived recreational value of lakes declines
when blooms form (Angradi et al., 2018), which in turn can negatively affect local economies
(Dodds et al., 2009). Despite the risk of economic loss, loss in biodiversity, and potential human
harm, the spatial dynamics of bloom development in spatially complex ecosystems such as
shallow lakes remain poorly characterized.

93 Shallow lakes have a large interface between the sediment and water relative to deeper 94 lakes, making them more susceptible to rapid changes in water residence time and nutrient inputs 95 (Christensen et al., 2013; Rennella & Quiros, 2006; Romo et al., 2013). Due to the expansive 96 littoral zones, shallow lakes can have large macrophyte beds which modify the light climate and 97 turbulence at the sediment-water interface (Andersen et al., 2017; Moller & Rordam, 1985; Vilas 98 et al., 2017). Many shallow lakes are also polymictic, experiencing multiple periods of 99 stratification followed by mixing during the ice-free season. During periods of water column 100 stability, some cyanobacteria taxa thrive, initiating blooms (Carey et al., 2012). Additionally, 101 episodic nutrient loading from the watershed during storm events (Carpenter et al., 2015; Kelly 102 et al., 2019), spatial gradients in nutrient availability due to stream inlets and morphology (e.g. 103 embayments), and wind-driven circulation (Schoen et al., 2014) can all contribute to spatial 104 heterogeneity of algal blooms over time in shallow lakes.

In order to better understand the spatial dynamics of algal blooms in shallow lakes, we performed intensive spatial sampling on Swan Lake (Iowa, USA), a spatially complex, shallow, hypereutrophic waterbody with a history of toxic cyanobacteria algal blooms. In addition to measuring algal pigments throughout the lake over the course of 122 days, we also measured temperature, dissolved oxygen, and pH. The spatial sampling captured two bloom events and

110	coincided with high frequency monitoring of the same variables using autonomous sensors
111	deployed at a fixed station (Ortiz et al. 2020). Using these data, we addressed the following
112	questions: 1) how does spatial variability of algae, dissolved oxygen, and pH change over the
113	course of a season, 2) are high frequency measurements at a fixed station an adequate
114	characterization of surface water dynamics in a shallow lake, and 3) what is the optimal spatial
115	sampling frequency to capture the mean state of a productive waterbody? Evaluating these
116	questions with data from a spatially complex, hypereutrophic lake will provide valuable
117	ecological and management-relevant insights into algal bloom dynamics.
118	
119	Methods
120	Study Site
121	Swan Lake (42.0396, -94.8454) has an average depth of 2 m, surface area of 40.5
122	hectares, and a shoreline development index value of 1.54. The watershed is 350 hectares with
123	92% of the land in agricultural use. The estimated water residence time is approximately 1.5
124	years. During the ice-free period of 2018, Swan Lake had an average total phosphorus
125	concentration of $280 \mu g L^{-1}$ and a total nitrogen concentration of 1.61 mg L^{-1} , making it
126	hypereutrophic (Carlson, 1977). Total nitrogen was measured as the sum of total Kjeldahl
127	nitrogen (method 351.2 v2, US EPA, 1993c) and nitrate + nitrite measured using the cadmium
128	reduction method (method 4500-NO3-F, US EPA, 1993a). Total phosphorus was measured using
129	the ascorbic acid method (method 365.1 v2, US EPA, 1993b). The average total alkalinity during
130	the same period was 139 mg CaCO ₃ L ⁻¹ determined through end point titration (APHA, 1998). In

131 addition to seasonal algal blooms, Swan Lake also has non-continuous beds of American lotus

132 (Nelumbo lutea) and sago pondweed (Stuckenia pectinata) that peak in biomass in the latter half

of the summer and then begin senescing. The main surface inlet to the lake enters on the western
side and the outlet is at the southern edge of the waterbody (Figure 1). There are no known
springs feeding the lake.

136

137 Field Methods

138 The spatial sampling occurred approximately weekly from day of year (DOY) 142 to 139 DOY 264, encompassing the late spring, summer, and early autumn. A total of 16 spatial 140 sampling events occurred over the course of the 122 days. Measurements of chlorophyll a, 141 phycocyanin, temperature, dissolved oxygen saturation and pH were taken 0.25 m below the 142 surface at 98 sampling stations using a YSI Pro DSS multiparameter sonde (Yellow Springs 143 Instrument, Yellow Springs, OH) suspended over the side of a 3-meter long jon boat equipped 144 with an outboard motor. The sensors, which included the fluorometric Total Algae (chlorophyll a 145 and phycocyanin), optical dissolved oxygen, and Ag/AgCl pH sensors, were calibrated weekly 146 prior to each sampling event according to manufacturer instructions. The sampling stations were 147 laid out in a 65 x 65 m grid across the lake (Figure 1) with each location measured in the same 148 order (north to south) for each sampling event. This spatial resolution was selected to allow for 149 many sampling locations to be measured in a relatively short window of time, thereby 150 minimizing the chance that the differences observed between sampling locations was not due to 151 time of day. Measurements were taken between 10:00 and 14:00, except for the first two and last 152 three weeks when sampling lasted until 16:00. Beginning on DOY 177 when submerged 153 macrophytes could be identified from the jon boat, the presence or absence of submerged or 154 floating leaf macrophytes was noted at each sampling station during each sampling event. These 155 weekly presence/absence data were used to construct the macrophyte distributions in Figure 1.

156 The fixed station high frequency monitoring of Swan Lake was performed using a YSI 157 EXO2 (Yellow Springs Instrument, Yellow Springs, OH) multiparameter sonde equipped with 158 the same sensors as the YSI ProDSS used for the spatial sampling. The sonde recorded 159 measurements of chlorophyll a, phycocyanin, dissolved oxygen saturation, and pH every 15 160 minutes. The instrument was deployed on DOY 135 over the deepest point in the lake (3.8 m 161 deep), hanging approximately 0.5m below the surface, and removed on DOY 264 after the 162 spatial sampling event on that day. The fixed station sonde was monitored weekly for drift and 163 calibrated according to manufacturer instructions when indicated by the quality control algorithm 164 in the KorEXO software. Hourly precipitation, wind speed, and wind direction were collected at 165 the Arthur N. Neu Airport in Carroll, Iowa, located 4.5 km from the lake, as a part of the 166 National Oceanic and Atmospheric Automated Surface Observatory System. The meteorological 167 data were used to aid in the interpretation of spatial dynamics over the course of the summer.

168

169 Data analysis

170 Spatial heterogeneity can be quantified by calculating the spatial variance (e.g., 171 coefficient of variation; CV) or spatial autocorrelation (Moran's I, Moran, 1950). Increasing 172 spatial variance is indicative of increasing patchiness in the ecosystem, such as areas of high-173 density algal biomass and areas of low-density biomass within a lake. Spatial autocorrelation 174 accounts for the location of those patches within the ecosystem in relationship to each other. 175 Local Moran's I quantifies how similar the abundance of algae is at one location compared to the 176 density of surrounding neighbors. When measured over time for variables that are indices of 177 algal biomass (e.g., the pigments chlorophyll a and phycocyanin), both of these metrics of spatial 178 heterogeneity can provide insight into the dynamics of algal bloom development. In models of

algal blooms, both spatial variance and autocorrelation are expected to be high during the bloomperiod (Buelo et al., 2018).

181 Spatial autocorrelation (AC) and the coefficient of variation (CV) were calculated for 182 each variable on each sampling date in order to evaluate the dynamics of these parameters over 183 time. Prior to analysis, extreme outliers in the algal pigments were removed from the spatial 184 dataset as they were well outside the operating range of the Total Algae sensor or there was 185 known interference with the sensor resulting in an inaccurate measurement. This resulted in five 186 chlorophyll a and three phycocyanin measurements being removed out of 3,136 total pigment 187 measurements. The spatial CV is the standard deviation of all of the spatial measurements for a 188 variable on a given sampling date divided by the mean of those measurements, expressed as a 189 percent. Spatial AC was calculated as the average value of local Moran's I with a queen's 190 distance neighbor list (92 meters) with equal weight (1/n) on neighbors, as to not impose any 191 assumptions on possible spatial patterns in the variables. We limited our analysis to surrounding 192 neighbors because distances beyond this have not shown high spatial autocorrelation of algal 193 pigments under experimental conditions (Butitta et al. 2017). Local Moran's I values near 1.0 194 reflect high spatial AC within neighbors, zero indicates a random distribution, whereas spatial 195 AC values nearing -1.0 indicate a perfectly dispersed distribution (e.g. checkerboard pattern) in 196 the variable being measured. As the spatial variability in temperature is mediated by physical 197 processes, we used the dynamics and extent of the spatial AC of temperature as a benchmark to 198 visually compare the dynamics of spatial AC in the other biological variables. This allowed us to 199 tease apart the effect of physically-versus biologically-driven spatial patterns. Additionally, in 200 order to better visualize the spatial patterns in chlorophyll a, phycocyanin, temperature,

dissolved oxygen, and pH over the course of the season, the data were interpolated using inverse
distance weighting across a 25m grid (Figure 2).

203 In order to evaluate if high frequency measurements at a fixed station are an adequate 204 characterization of the surface water dynamics in a shallow lake, we compared the measurements 205 taken by the fixed station sonde during the same time period as a spatial sampling event. High 206 frequency data from the fixed station sonde was trimmed to the period that we sampled the lake 207 spatially. A t-test with a Bonferroni correction, to account for the multiple comparisons, was 208 performed to compare the distribution from the 98 sampling stations to the fixed station 209 measurements from the same day for each of the four biologically-mediated variables, 210 chlorophyll, phycocyanin, dissolved oxygen, and pH. In addition to comparing fixed sonde 211 values to the spatial sampling, we also used the spatial data to identify locations in the lake that 212 were consistently representative of mean conditions, and therefore ideal locations for fixed 213 station monitoring. We identified locations in the lake for each sampling event that had 214 measurements within the range of \pm one standard deviation from the mean for each biologically 215 mediated variable (all variables except temperature). We then collated these locations across all 216 sampling dates to identify which of the 98 sampling locations had measurements that most 217 consistently represented the mean conditions of the lake.

Finally, we performed a rarefaction analysis to evaluate the optimal spatial sampling frequency to capture the mean value of the biologically-mediated variables. This was done by randomly selecting *n* number of spatial sampling data points (n=2-97) during a sampling event, calculating the mean value from that subset, and then calculating the root mean square error (RMSE), comparing the mean estimate from the subset to the mean of all sampling points during that event. This calculation was repeated 1000 times for each value of *n*, and each iteration was

224	then averaged. The averaged RMSE values for each subset of n were fit using a local polynomial
225	regression with a smoothing factor of 0.1 and each sampling event's RMSE curve was
226	standardized by subtracting the mean of all iterations ("global mean") from the mean at n
227	number of stations, to aid in visual comparison. The spatial data are available through (Ortiz &
228	Wilkinson, 2019) and the fixed station data are available in (Ortiz et al., 2019) and further
229	analyzed in Ortiz et al. (2020). All analyses were performed using R 4.0.2 (R Core Team, 2020)
230	using the gstat (Pebesma, 2004), rstatix (Kassambara, 2020), and sf packages (Pebesma, 2018).
231	
232	Results

233 There were two bloom events during the summer of 2018 in Swan Lake. The first bloom 234 occurred from DOY 156 – 184 and was dominated by the diatom Aulacoseira spp. based on a 235 sample taken on DOY 177 examined under a compound microscope at 400x magnification. The 236 phycocyanin concentrations on DOY 177 were the lowest during this first bloom period (Figure 237 2), and no cyanobacteria were identified in the sample. The second bloom, peaking on DOY 236, 238 was dominated by the cyanobacterium *Microcystis spp*. There were also two large precipitation 239 events during the summer, occurring after sampling on DOY 170 and lasting through DOY 171, 240 and on DOY 232 (Figure 2; Supplemental Figure 1). The maximum wind speed recorded during 241 the first precipitation event was 10.8 m s⁻¹ coming from the southwest and 11.8 m s⁻¹ during the 242 second precipitation event coming from the southeast. During the first half of the summer (DOY 243 142 - 191) the prevailing winds 24 hours prior to the sampling events were from the south, 244 switched to being predominantly from the north from DOY 198 - 219, and then varied in 245 direction for the rest of the season (Figure 2). The median wind speed for the first period when 246 winds were out of the south was 3.6 m s⁻¹ (Figure 3b). When the winds switched to being

predominantly from the north between DOY 198 - 219, the median wind speed was lower at 2.5 m s⁻¹ (Figure S1).

249 Spatial dynamics

250 During the two bloom periods there was not a latitudinal or longitudinal trend in 251 chlorophyll *a* concentrations; instead, there were patches of high chlorophyll *a* concentration on 252 otherwise low-concentration dates (Figure 2). Unlike chlorophyll a, phycocyanin had a strong 253 latitudinal trend with higher concentrations in the northern portion (sample sites A1-G4) of the 254 lake during the first bloom. This spatial pattern is readily observed on DOY 184 but is also 255 noticeable for many of the sampling events during the first bloom (Figure 2). During sampling 256 events with a strong latitudinal gradient in phycocyanin (DOY 166 - 184, and 236) the mean 257 concentration in the northern portion of the lake was nearly double the concentration in the 258 southern portion of the lake (7.29 and $3.76 \,\mu g \, L^{-1}$, respectively). On these dates, the prevailing 259 winds 24 hours prior to the sampling event were out of the north (Supplemental Figure S1), yet 260 the lowest concentrations of phycocyanin were found in the southern portion of the lake. Even 261 when the lake was not blooming, there were patches of high concentrations of phycocyanin in 262 the northern portion of the lake (e.g., DOY 212), located among the densest patch of American 263 Lotus (Figure 1). The average phycocyanin concentrations at the sampling locations within the 264 American Lotus patch was higher than the average concentration in the rest of the lake for 14 of 265 the 16 sampling events (Figure S2).

The daytime saturation of dissolved oxygen varied the most out of the five variables monitored, ranging from borderline hypoxic (30% saturation) to supersaturated (up to 350%) (Figure 2). While the dissolved oxygen saturation increased near the peak of the bloom, the highest average saturation was on DOY 191, after the first bloom had collapsed. There was a

270 weak pattern over the course of the season of higher saturation in the northern portion of the 271 lake, similar to the distribution of higher phycocyanin concentrations. However, within the 272 northern portion of the lake, regions of low dissolved oxygen saturation formed in the surface 273 waters, particularly later in the summer (Figure 1). Beginning on DOY 198, the mean dissolved 274 oxygen concentration in the American lotus patch was consistently lower than the average for the 275 rest of the lake until DOY 250 (Figure S2). The distribution of pH also had a weak spatial pattern 276 during the summer, with slightly elevated values in the northern portion of the lake during the 277 first bloom (e.g. DOY 177; Figure 2). While pH was elevated at the onset of the first bloom 278 period from DOY 149 – 170, it was highest overall on DOY 191 and 198 after the first bloom 279 had collapsed. Unlike the other variables, temperature had a subtle south to north latitudinal 280 gradient with warmer temperatures in the southern portion of the lake and colder in the north 281 during the latter half of the summer (Figure 2). On average this difference between the northern 282 portion of the lake and the southern was 0.5°C. The warmest day of sampling was DOY 191. 283 Spatial variability in algal pigments during the first bloom event was low, with two 284 exceptions. There was an increase in the CV of chlorophyll a on the last day of the bloom (DOY 285 184; Figure 3a) that continued to increase as the bloom collapsed. There was also a temporary 286 increase in phycocyanin CV during the first bloom on DOY 177 (Figure 3b), coinciding with a 287 temporary decline in phycocyanin concentration across the lake. The CV of both algal pigments 288 was higher than the CV of temperature over the course of the entire sampling period. 289 Conversely, the CV of pH and dissolved oxygen were elevated during the first bloom 290 period, with pHCV declining and remaining low after the first bloom (Figure 3c) and dissolved 291 oxygen CV only temporarily declining after the first bloom (Figure 3d). Temperature had low 292 variability throughout the first bloom until DOY 177, when the lake began to heat up, peaking in

both temperature and spatial variability on DOY 191 (Figure 3e). Between the first and second
bloom, DOY 191-226, there was a decrease in spatial variability among the algal pigments and
pH as the bloom collapsed, while temperature and dissolved oxygen CV remained relatively high
and variable. During the second bloom period, CV was low for all variables except for
chlorophyll *a*. In general, the CV, of temperature and pH, expressed as a percentage, was an
order of magnitude lower than the other variables.

299 Spatial autocorrelation (AC), quantified as local Moran's I, did not fall substantially 300 below 0 for any of the variables and peaked at 0.79 among all variables (Figure 3). The highest 301 AC value for chlorophyll a and phycocyanin was during the first bloom event (Figure 3f, g); 302 however, phycocyanin AC also increased substantially during the second bloom. During the first 303 bloom, the AC of temperature varied similarly to both pigments' AC, particularly phycocyanin, 304 but became decoupled after the bloom collapsed. While the AC of temperature remained high 305 during the inter-bloom period, the AC of the pigments was substantially lower. Conversely, the 306 dynamics of AC of temperature, dissolved oxygen and pH remained coupled throughout the 307 summer (Figure 3h, i). Dissolved oxygen saturation and pH both increased in AC during the first 308 bloom and then declined throughout the rest of the season with the exception of a minor increase 309 in AC during the second bloom event.

310

311 Fixed station versus spatial sampling

There were a greater number of days with a significant difference between the spatial and fixed station measurements than days in which the data sets were not significantly different (Figure 4). Among all 64 comparisons (4 variables × 16 sampling events), the spatial and fixed station data sets had a means that were not significantly different 37.5% of the time.

316 Phycocyanin had the greatest number of events with similar values, with 7 of the 16 sampling 317 events having non-statistically different mean values measured spatially and at the fixed station 318 (Figure 4b). These occurrences were mainly during non-bloom periods. However, even when the 319 mean phycocyanin values were similar between the sampling methods on a given day, the range 320 of values captured by the fixed station was five times less than the variability captured in the 321 spatial data. This pattern of infrequent occurrences of similar mean values between the two 322 methods during non-bloom periods and a diminished range in the fixed station data, was shared 323 to a degree, among the other three variables as well. Interestingly, dissolved oxygen saturation 324 only had 5 out of the 16 events with means that were not significantly different, all of which 325 occurred when the lake was above 100% saturation (Figure 4c).

While a majority of the comparisons between the fixed station and spatial data indicate that the algal pigments had a larger range of values in the spatial data, there were a handful of instances where the opposite was true. During the first bloom, the fixed station sonde measured a wide range of chlorophyll *a* concentrations and had a higher mean chlorophyll *a* for all dates (Figure 4a). Similarly, we observed higher mean phycocyanin at the fixed station sonde on DOY 156, 166, 177, 191, and 219 (Figure 4b). However, this pattern did not hold true for dissolved oxygen or pH (Figure 5c, d).

The spatial sampling sites that most consistently captured the mean values in the lake on a given sampling date were in the northwest portion of the lake, near the inlet. The best performing site for all variables was site E3, just west of the American lotus patch and adjacent to a bed of sago pondweed (Figure 1). The four biologically-mediated variables from sample site E3 were within the mean (± standard deviation) range of all of the spatial measurements 95% of the time. The second best performing location was in the middle of the American lotus patch, site

D4, with the values from this site being within the mean (± standard deviation) range 92% of the
time. The site where the fixed station was located, site H2, was only within the mean (± standard
deviation) range 58% of the time.

342

343 Optimal Spatial Resolution

344 In order to evaluate the spatial sampling resolution needed to capture the mean state of 345 the surface water on a given day, we performed a rarefaction analysis for each variable and each 346 sampling event, calculating the root mean squared error (RMSE) of a subset of sampling 347 locations compared to the mean value of all 98 measurements that day. The plateaus of the 348 RMSE curves from the rarefaction analysis were used to evaluate the smallest number of spatial 349 sampling locations needed to capture the mean across the lake during that sampling event (Figure 350 5). Additionally, we also evaluated the temporal pattern of the minimum number of sampling 351 locations needed to capture the mean.

352 Mean values were underestimated for all variables on all sampling dates when there were 353 less than 10 sampling stations (Figure 5). However, the severity of the underestimation differed 354 among the variables. The rarefaction analysis for chlorophyll a indicated that 10 - 30 sampling 355 locations was sufficient for capturing the mean chlorophyll *a* in Swan Lake, otherwise the mean 356 concentration would be under estimated (Figure 5a). When an algal bloom was occurring it took 357 more sampling locations to near the mean chlorophyll a concentration on that date. However, 358 when the bloom was particularly patchy during development (DOY 226) or collapse (DOY 191), 359 including a larger number of sampling locations led to overestimating the mean chlorophyll a 360 concentration as locations with high concentrations were over-represented in the data set. There 361 were similar patterns in phycocyanin RMSE with most sampling dates plateauing between 20 -

362 30 sampling locations with a few exceptions (Figure 5b). For DOYs 156-170 (rise of the first 363 bloom) and 212, at least 60 sampling locations were needed to capture the overall mean in 364 phycocyanin for that sampling date. Dissolved oxygen saturation and pH were generally well 365 characterized by approximately 10 - 15 sampling locations as both had a majority of dates in 366 which the RMSE curves plateaued at that spatial sampling resolution (Figure 5c, d). However, at 367 the beginning (DOY 154), peak (DOY 184), and end (DOY 205) of the first bloom, twice as 368 many sampling locations were needed to capture the mean dissolved oxygen. Only two dates 369 required more sampling locations for pH to capture the mean, DOY 177 and 198, which 370 plateaued at approximately 40 sampling locations. The largest RMSE were observed during 371 bloom conditions for all variables: DOY 177 had the largest error for phycocyanin and pH, while 372 the largest RMSE was on DOY 184 for dissolved oxygen and on DOY 236 for chlorophyll a 373 (Figure 5).

374

375 Discussion

376 The spatial heterogeneity of water quality parameters was highly dynamic in Swan Lake, 377 a shallow, hypereutrophic, temperate waterbody. The temporal dynamics in heterogeneity were 378 driven in part by the two blooms, the peaks of which were preceded by large precipitation events. 379 These rain events could have delivered nutrients from the agriculturally dominated watershed 380 into the lake from the northern inlet, helping to fuel the subsequent algal blooms and the spatial 381 patterns observed during blooms (Stockwell et al. 2020). However, there are also a number of 382 other factors that likely contributed to the spatial variability and pattern during and following 383 these bloom events, including the prevailing wind direction prior to sampling, the bathymetry of

the basin and location of the surface inlet, and the potential for dense macrophyte beds tocontribute to local patchiness.

386 The spatial patterns that the algal blooms created were consistent with the expectations 387 from previous modeling and experimental work that spatial AC increases as algal blooms 388 develop (Buelo et al., 2018; Butitta et al., 2017; Serizawa et al., 2008). This pattern was the 389 strongest for phycocyanin, evident by the strong latitudinal gradient in concentrations during the 390 bloom periods. The sampling dates with phycocyanin concentration gradients (e.g., DOY 166, 391 177, 184, 236) coincided with persistent winds from the south 24 hours prior to the sampling 392 event, which likely resulted in the higher concentration of algal cells in the northern portion of 393 the lake. The effect of persistent wind directions influencing the distribution of cyanobacteria has 394 also been documented in other shallow eutrophic lakes (Wu et al. 2010). The shallow sediments 395 of the northern basin were also likely a source of akinete recruitment (Karlsson-Elfgren and 396 Brunburg, 2004), further contributing to the higher concentrations of phycocyanin in the northern 397 portion of the lake during the first bloom. Augmented nutrient availability in the northern part of 398 the lake due to external loading from the watershed through the surface inlet and internal loading 399 from the sediments overlain by an unstratified water column (Song and Burgin 2017) may have 400 further amplified the phytoplankton gradient, particularly following precipitation events. Finally, 401 the tendency of the dominant cyanobacteria taxa *Microcystis spp.* to form surface scums likely 402 enhanced the spatial patterns observed with our surface sampling approach.

The sampling dates with a strong gradient of phytoplankton concentrations from north to south also resulted in north-south gradients in water chemistry. On these dates, both dissolved oxygen and pH formed a gradient of high values in the northern portion of the lake and lower values in the south, which would be expected with greater primary production where

407 phytoplankton concentrations were highest. The spatial patterns in the surface water chemistry 408 demonstrate how phytoplankton spatial distribution, driven by wind, can create hot spots and 409 moments of biogeochemical activity within lakes (McClain et al. 2003) that may be missed with 410 traditional, single-station sampling. The dense patch of floating leaf American lotus macrophytes 411 also created a hot spot of biogeochemical activity.

412 Macrophyte beds can have a large local influence on water chemistry by inducing 413 stratification, decreasing flow and trapping particles, and modifying the light environment 414 (Green, 2006; Vilas et al., 2017). For 87.5% of the season the phycocyanin concentrations were 415 higher in the bed of American lotus than concentrations elsewhere in the lake. In fact, even on 416 sampling dates when phycocyanin concentrations were otherwise low (e.g., DOY 212), the 417 American lotus patch can be identified based on the phycocyanin concentrations that are nearly 418 twice as high as the rest of the lake. We hypothesize that the macrophyte patch allowed for 419 microstratification in the water column and reduced wind-driven flow. These physical conditions 420 are likely to favor cyanobacteria dominance and the formation of surface scums. Similarly, the 421 dissolved oxygen concentrations in the American lotus patch became consistently lower than the 422 rest of the lake later in the summer, likely due to the plants beginning to senesce, creating a hot 423 spot of decomposition, decreasing both dissolved oxygen and pH (Vilas et al., 2017). While there 424 isn't strong evidence in the data that the other submerged macrophyte beds had a similarly strong 425 effect on water chemistry, the data from the American lotus patch illustrates how macrophytes 426 can contribute to local patchiness and overall spatial heterogeneity.

427

428 Considerations for Monitoring

429 The variables that we measured in this study are often the target of water quality 430 monitoring as the dynamics of these variables coincide with changes in ecosystem function and 431 services. Monitoring is often performed at a fixed station over time to capture the dynamics of 432 the ecosystem, but this strategy could potentially result in missed information about the 433 ecosystem's behavior. While the temporal dynamics of all the variables were synchronous 434 between the fixed station and spatial sampling data sets in Swan Lake, our conclusions regarding 435 the magnitude of the blooms and variability in the lake's structure would have been substantially 436 different relying solely on the fixed station data. Among the four biologically-mediated 437 variables, only 37.5% of the fixed station estimates of the mean state of the lake statistically 438 matched the estimate from the spatial sampling. The vast majority of those instances (96%) 439 occurred during non-bloom periods, which also coincided with lower wind speed conditions, no 440 prevailing wind direction, and no major precipitation events. The large difference between the 441 spatial sampling and fixed station measurements of algal pigments during blooms was likely 442 driven, in part, by the depth of the sensors at the fixed station and the variable accumulation of 443 cyanobacteria at the surface of the lake dependent upon environmental conditions and the 444 dominant taxa (Chaffin et al., 2020). It is clear from our data that during periods of heightened 445 biological activity such as blooms, fixed station monitoring is unlikely to be representative of the 446 mean ecosystem state in shallow lakes.

Despite the high degree of horizontal spatial variability that has been documented in this study and others (Loken et al. 2019, Van de Bogert et al. 2012, Buttita et al. 2017), fixed station designs are widely used in water quality monitoring programs. In Swan Lake, we determined that the historical location for water quality monitoring, where the fixed station sensors were deployed, was one of the least-representative locations for mean conditions in the lake. Given the

452 hypereutrophic state of the lake, the most immediate management concerns are toxic 453 cyanobacteria blooms and summer fish kills due to low dissolved oxygen. Yet, the mean value of 454 these variables (phycocyanin and dissolved oxygen) across the lake were only captured by the 455 fixed station sensors 58% of the time. While selecting a fixed station site for high frequency 456 sensor deployment includes many considerations including the location of previous data 457 collection and management needs, based on our analysis we would advise performing a spatial 458 survey to identify if and when the fixed station site is representative of mean conditions in the 459 lake. A complementary spatial survey will help contextualize the fixed station dynamics and 460 provide additional, management-relevant information about the lake. It's also important to consider the trade-offs between high frequency fixed station 461 462 monitoring and higher resolution, but less frequent spatial monitoring. High frequency 463 monitoring at a single station provides insight into ecosystem function such as metabolism 464 (Staehr et al., 2012), early warnings of impending regime shifts (Carpenter et al., 2011; 465 Wilkinson et al., 2018), and crucial information on diel variability in limnological conditions 466 (Andersen et al., 2017). However, as we observed in Swan Lake, the spatial variability within a 467 given day often exceeds the temporal variability at a single point in a shallow lake. Without the 468 spatial sampling snapshots, we would have underestimated the magnitude of the algal blooms, 469 hampering our limnological understanding of the ecosystem's functioning and impeding our 470 ability to accurately estimate rates such as methane emissions on a global scale (DelSontro et al. 471 2018).

From a practical stand point, the understanding gleaned from the spatial sampling could help managers design targeted algal toxin monitoring or management interventions to help control fish habitat quality in persistently hypoxic areas (Bardshaw et al., 2015). However, the

475 time and cost investment in repeated spatial sampling at the resolution performed in this study 476 may not be feasible for both research and management programs. The rarefaction analysis we 477 performed for all four of the key water quality monitoring variables revealed that minimal 478 information was gained after ~ 30 locations were sampled across many conditions and variables. 479 Often 12-20 sample locations across the 40.5 ha lake (or a 1-2 samples per hectare) was 480 sufficient to capture the spatial variability within the lake, with a few exceptions. These 481 exceptions occurred during times of higher variability such as when the blooms were just starting 482 or when the bloom began to collapse. The need for a higher spatial resolution during bloom 483 events to fully capture their variability has also been found using remote sensing techniques in 484 other, larger lakes (Lekki et al., 2019). As the spatial resolution of remote sensing technologies 485 continues to improve, it may become more cost effective to capture the spatial heterogeneity of 486 algal pigments in small lakes over time. However, one of the benefits of manual spatial sampling 487 is being able to pair other measurements such as dissolved oxygen, pH, and nutrients (e.g., 488 nitrate; Loken et al., 2018; Pellerin et al., 2016) with information on the distribution of algal 489 biomass.

490 Our intensive spatial monitoring of a shallow, hypereutrophic lake revealed how spatially 491 heterogeneous shallow lakes are over the course of a single season and allowed us to tease apart 492 the drivers of that spatial heterogeneity. We found that variability was greatest during 493 biologically-intensive periods, such as during algal blooms and in dense floating-leaf macrophyte 494 beds, and that failure to capture this variability would have hampered our understanding of the 495 ecosystem's functioning and overall mean state. Small lakes such as Swan Lake dominate the 496 global distribution of waterbodies (Verpoorter et al. 2014). Adequately capturing and 497 characterizing the magnitude of variability in production of these waterbodies is important given

498	their role in mediating global nutrient cycles (Downing et al. 2010, Biddanda et al. 2017),
499	especially methane emissions (DelSontro et al. 2018, Loken et al. 2019). Our data provided an
500	estimate of the spatial resolution needed to capture the dynamics in ecosystems similar to Swan
501	Lake and a method which could be readily adapted to other ecosystems. While our results
502	provide new understanding of the magnitude and temporal dynamics of spatial heterogeneity in
503	shallow lakes, continued investigation of horizontal spatial heterogeneity in a range of aquatic
504	ecosystems, from oligotrophic to eutrophic, is needed to better understand the structure and
505	drivers of horizontal spatial variability in lakes.
506	
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515	References
516 517 518 519 520	 American Public Health Association (APHA), American Water Works Association (AWWA), and the Water Environmental Federation (WEF). 1998. Standard Methods for Examinations of Water and Wastewater, 20th ed. United Book Press, Inc. Baltimore, Maryland. Andersen, M. R., Kragh, T., & Sand-Jensen, K. (2017). Extreme diel dissolved oxygen and
521 522 523 524 525	 carbon cycles in shallow vegetated lakes. <i>Proceedings of the Royal Society B-Biological Sciences</i>, 284(1862), doi:10.1098/rspb.2017.1427 Angradi, T. R., Ringold, P. L., & Hall, K. (2018). Water clarity measures as indicators of recreational benefits provided by US lakes: Swimming and aesthetics. <i>Ecological Indicators</i>, 93, 1005-1019. doi:10.1016/j.ecolind.2018.06.001

- Bardshaw, E. L., Allen, M. S., & Netherland, M. (2015). Spatial and temporal occurrence of
 hypoxia influences fish habitat quality in dense Hydrilla verticillata. *Journal of Freshwater Ecology*, 30(4), 491-502.
- Biddanda, B. A. (2017). Global significance of the changing freshwater carbon cycle. *EOS*, 98,
 doi:10.1029/2017EO069751
- Boehrer, B., & Schultze, M. (2008). Stratification of lakes. *Reviews of Geophysics*, 46(2),
 doi:10.1029/2006rg000210
- Buelo, C. D., Carpenter, S. R., & Pace, M. L. (2018). A modeling analysis of spatial statistical
 indicators of thresholds for algal blooms. *Limnology and Oceanography Letters*, 3(5),
 384-392. doi:10.1002/lol2.10091
- Butitta, V. L., Carpenter, S. R., Loken, L. C., Pace, M. L., & Stanley, E. H. (2017). Spatial early
 warning signals in a lake manipulation. *Ecosphere*, 8(10), doi:10.1002/ecs2.1941
- 538 Carlson, R. E. (1977). Trophic State Index for Lakes. *Limnology and Oceanography*, 22(2), 361 539 369. doi:10.4319/lo.1977.22.2.0361
- Carpenter, S., Booth, E., Kucharik, C., & Lathrop, R. (2015). Extreme daily loads: role in annual
 phosphorus input to a north temperate lake. *Aquatic Sciences*, 77(1), 71-79.
 doi:10.1007/s00027-014-0364-5
- 543 Carpenter, S. R., Arani, B. M. S., Hanson, P. C., Scheffer, M., Stanley, E. H., & Van Nes, E.
 544 (2020). Stochastic dynamics of Cyanobacteria in long-term high-frequency observations
 545 of a eutrophic lake. *Limnology and Oceanography Letters*, 5(5), 331-336.
 546 doi:10.1002/lol2.10152
- 547 Carpenter, S. R., Cole, J. J., Pace, M. L., Batt, R., Brock, W. A., Cline, T., . . . Weidel, B. (2011).
 548 Early Warnings of Regime Shifts: A Whole-Ecosystem Experiment. *Science*, *332*(6033),
 549 1079-1082. doi:10.1126/science.1203672
- Chaffin, J. D., Kane, D. D., & Johnson, A. (2020). Effectiveness of a fixed-depth sensor
 deployed from a buoy to estimate water-column cyanobacterial biomass depends on wind
 speed. *Journal of Environmental Sciences*, 93, 23-29, doi:10.1016/j.jes.2020.03.003
- Christensen, J. P. A., Sand-Jensen, K., & Staehr, P. A. (2013). Fluctuating water levels control
 water chemistry and metabolism of a charophyte-dominated pond. *Freshwater Biology*,
 555 58(7), 1353-1365. doi:10.1111/fwb.12132
- Codd, G. A., Morrison, L. F., & Metcalf, J. S. (2005). Cyanobacterial toxins: risk management
 for health protection. *Toxicology and Applied Pharmacology*, 203(3), 264-272.
 doi:10.1016/j.taap.2004.02.016
- Corbel, S., Mougin, C., & Bouaicha, N. (2014). Cyanobacterial toxins: Modes of actions, fate in
 aquatic and soil ecosystems, phytotoxicity and bioaccumulation in agricultural crops.
 Chemosphere, 96, 1-15. doi:10.1016/j.chemosphere.2013.07.056
- 562 Cotterill, V., Hamilton, D. P., Puddick, J., Suren, A., & Wood, S. A. (2019). Phycocyanin
 563 sensors as an early warning system for cyanobacteria blooms concentrations: a case study
 564 in the Rotorua lakes. *New Zealand Journal of Marine and Freshwater Research*, 53(4),
 565 555-570. doi:10.1080/00288330.2019.1617322
- Davis, T. W., Berry, D. L., Boyer, G. L., & Gobler, C. J. (2009). The effects of temperature and
 nutrients on the growth and dynamics of toxic and non-toxic strains of Microcystis during
 cyanobacteria blooms. *Harmful Algae*, 8(5), 715-725. doi:10.1016/j.hal.2009.02.004
- DelSontro, T., Beaulieu, J. J., & Downing, J. A. (2018). Greenhouse gas emissions from lakes
 and impoundments: Upscaling in the face of global change. *Limnology and Oceanography Letters*, 3(3), 64-75. doi:10.1002/lol2.10073

572 Dodds, W. K., Bouska, W. W., Eitzmann, J. L., Pilger, T. J., Pitts, K. L., Riley, A. J., . . . 573 Thornbrugh, D. J. (2009). Eutrophication of US Freshwaters: Analysis of Potential 574 Economic Damages. Environmental Science & Technology, 43(1), 12-19. 575 doi:10.1021/es801217g 576 Downing, J. A. 2010. Emerging global role of small lakes and ponds: little things mean a lot. 577 *Limnetica* 29(1), 9-24. 578 George, D. G., & Heaney, S. I. (1978). Factors influencing spatial-distribution of phytoplankton 579 in a small productive lake. Journal of Ecology, 66(1), 133-155. doi:10.2307/2259185 580 Gilbert, P. M. (2017). Eutrophication, harmful algae and biodiversity - Challenging paradigms in 581 a world of complex nutrient changes. Marine Pollution Bulletin, 124(2), 591-606. 582 doi:10.1016/j.marpolbul.2017.04.027 583 Green, J. C. (2006). Effect of macrophyte spatial variability on channel resistance. Advances in 584 Water Resources, 29(3), 426-438. doi:10.1016/j.advwatres.2005.05.010 585 Hansen, A. M., Andersen, F. O., & Jensen, H. S. (1997). Seasonal pattern in nutrient limitation 586 and grazing control of the phytoplankton community in a non-stratified lake. Freshwater 587 *Biology*, *37*(3), 523-534. doi:10.1046/j.1365-2427.1997.00182.x 588 Kassambara, A. (2020). rstatix: Pipe-Friendly Framework for Basic Statistical Tests. R package 589 version 0.6.0. Retrieved from https://cran.r-project.org/package=rstatix 590 Karlsson-Elfgren, I., & Brunberg, A. K. (2004). The importance of shallow sediments in the 591 recruitment of Anabaena and Aphanizomenon (Cyanophyceae). Journal of Phycology, 592 40(5), 831-836. doi:10.1111/j.1529-8817.2004.04070.x 593 Kelly, P. T., Renwick, W. H., Knoll, L., & Vanni, M. J. (2019). Stream Nitrogen and Phosphorus 594 Loads Are Differentially Affected by Storm Events and the Difference May Be 595 Exacerbated by Conservation Tillage. Environmental Science & Technology, 53(10), 596 5613-5621. doi:10.1021/acs.est.8b05152 597 Laas, A., Noges, P., Koiv, T., & Noges, T. (2012). High-frequency metabolism study in a large 598 and shallow temperate lake reveals seasonal switching between net autotrophy and net 599 heterotrophy. Hydrobiologia, 694(1), 57-74. doi:10.1007/s10750-012-1131-z 600 Landsberg, J. H. (2002). The effects of harmful algal blooms on aquatic organisms. Reviews in 601 Fisheries Science, 10(2), 113-390. doi:10.1080/20026491051695 602 Lekki, J., Deutsch, E., Sayers, M., Bosse, K., Anderson, R., Tokars, R., & Sawtell, R. (2019). 603 Determining remote sensing spatial resolution requirements for the monitoring of harmful 604 algal blooms in the Great Lakes. Journal of Great Lakes Research, 45(3), 434-443. 605 doi:10.1016/j.jglr.2019.03.014 606 Loken, L. C., Crawford, J. T., Dornblaser, M. M., Striegl, R. G., Houser, J. N., Turner, P. A., & 607 Stanley, E. H. (2018). Limited nitrate retention capacity in the Upper Mississippi River. 608 Environmental Research Letters, 13(7). doi:10.1088/1748-9326/aacd51 609 Loken, L. C., Crawford, J. T., Schramm, P. J., Stadler, P., Desai, A. R. & Stanley, E. H. (2019). 610 Large Spatial and Temporal Variability of Carbon Dioxide and Methane in a Eutrophic 611 Lake. Journal of Geophysical Research-Biogeosciences, 124(7), 2248-2266. 612 10.1029/2019jg005186 613 McClain, M. E., Boyer, E. W., Dent, C. L., Gergel, S. E., Grimm, N. B., Groffman, P. M., ... 614 Pinay, G. (2003). Biogeochemical hot spots and hot moments at the interface of terrestrial and aquatic ecosystems. *Ecosystems*, 6(4), 301-312. doi:10.1007/s10021-003-0161-9 615 616 Medeiros, A. S., Biastoch, R. G., Luszczek, C. E., Wang, X. A., Muir, D. C. G., & Quinlan, R. 617 (2012). Patterns in the limnology of lakes and ponds across multiple local and regional

618 environmental gradients in the eastern Canadian Arctic. Inland Waters, 2(2), 59-76. 619 doi:10.5268/iw-2.2.427 620 Moller, T. R., & Rordam, C. P. (1985). Species numbers of vascular plants in relation to area, 621 isolation and age of ponds in Denmark. Oikos, 45(1), 8-16. doi:10.2307/3565216 622 Moran, P. A. P. (1950). Notes on Continuous Stochastic Phenomena. Biometrika, 37(1-2), 17-23. 623 doi:10.2307/2332142 624 Moreno-Ostos, E., Cruz-Pizarro, L., Basanta, A., & George, D. G. (2009). Spatial Heterogeneity 625 of Cyanobacteria and Diatoms in a Thermally Stratified Canyon-Shaped Reservoir. International Review of Hydrobiology, 94(3), 245-257. doi:10.1002/iroh.200811123 626 627 Ortiz, D., Palmer, J., & Wilkinson, G. (2019). Hypereutrophic lake sensor data during summer 628 algae blooms in Iowa, USA, 2014 - 2018 ver 1. Environmental Data Initiative. 629 doi:10.6073/pasta/30070d41fbcdf36387f33d9108f570f8 Ortiz, D., & Wilkinson, G. (2019). Hypereutrophic lake spatial sensor data during summer 630 631 bloom, Swan Lake, Iowa, USA 2018 ver 1. Environmental Data Initiative. 632 doi:10.6073/pasta/2c0ca177438a3d422925811514e86cd8 Ortiz, D., Palmer, J., & Wilkinson, G. (2020). Detecting changes in statistical indicators of 633 634 resilience prior to algal blooms in shallow eutrophic lakes. *Ecosphere*, 11(10). 635 doi:10.1002/ecs2.3200 636 Pace, M. L., Batt, R. D., Buelo, C. D., Carpenter, S. R., Cole, J. J., Kurtzweil, J. T., & Wilkinson, 637 G. M. (2017). Reversal of a cyanobacterial bloom in response to early warnings. Proceedings of the National Academy of Sciences of the United States of America, 638 639 114(2), 352-357. doi:10.1073/pnas.1612424114 640 Pebesma, E. J. (2018). Simple Features for R: Standardized Support for Spatial Vector Data. R 641 Journal, 10(1), 439-446. 642 Pebesma, E. J. (2004). Multivariable geostatistics in S: the gstat package. Computers & 643 *Geosciences*, 30(7), 683-691. doi:10.1016/j.cageo.2004.03.012 644 Pellerin, B. A., Stauffer, B. A., Young, D. A., Sullivan, D. J., Bricker, S. B., Walbridge, M. R., 645 ... Shaw, D. M. (2016). Emerging tools for continuous nutrient monitoring networks: 646 sensors advancing science and water resources protection. Journal of the American Water 647 Resources Association, 52(4), 993-1008. doi:10.1111/1752-1688.12386 648 R Core Team. (2020). R: A language and environment for statistical computing. R Foundation 649 for Statistical Computing, Vienna, Austria. Retrieved from https://www.R-project.org/ 650 Read, J. S., Hamilton, D. P., Jones, I. D., Muraoka, K., Winslow, L. A., Kroiss, R., ... Gaiser, E. 651 (2011). Derivation of lake mixing and stratification indices from high-resolution lake 652 buoy data. Environmental Modelling & Software, 26(11), 1325-1336. 653 doi:10.1016/i.envsoft.2011.05.006 654 Rennella, A. M., & Quiros, R. (2006). The effects of hydrology on plankton biomass in shallow 655 lakes of the Pampa Plain. Hydrobiologia, 556, 181-191. doi:10.1007/s10750-005-0318-y 656 Romo, S., Soria, J., Fernandez, F., Ouahid, Y., & Baron-Sola, A. (2013). Water residence time 657 and the dynamics of toxic cyanobacteria. Freshwater Biology, 58(3), 513-522. 658 doi:10.1111/j.1365-2427.2012.02734.x 659 Rychtecky, P., & Znachor, P. (2011). Spatial heterogeneity and seasonal succession of 660 phytoplankton along the longitudinal gradient in a eutrophic reservoir. Hydrobiologia, 663(1), 175-186. doi:10.1007/s10750-010-0571-6 661 662 Schilder, J., Bastviken, D., van Hardenbroek, M., Kankaala, P., Rinta, P., Stotter, T., & Heiri, O. 663 (2013). Spatial heterogeneity and lake morphology affect diffusive greenhouse gas

664 emission estimates of lakes. Geophysical Research Letters, 40(21), 5752-5756. 665 doi:10.1002/2013gl057669 666 Schoen, J. H., Stretch, D. D., & Tirok, K. (2014). Wind-driven circulation patterns in a shallow 667 estuarine lake: St Lucia, South Africa. Estuarine Coastal and Shelf Science, 146, 49-59. doi:10.1016/j.ecss.2014.05.007 668 669 Serizawa, H., Amemiya, T., & Itoh, K. (2008). Patchiness in a minimal nutrient - phytoplankton 670 model. Journal of Biosciences, 33(3), 391-403. doi:10.1007/s12038-008-0059-y 671 Smith, C., D. (2018). Temporal and Spatial Monitoring of Cyanobacterial Blooms at Willow 672 Creek Reservoir, North-Central Oregon. Retrieved from U.S. Geological Survey 673 Scientific Investigations Report: 674 Solomon, C. T., Bruesewitz, D. A., Richardson, D. C., Rose, K. C., Van de Bogert, M. C., 675 Hanson, P. C., ... Zhu, G. W. (2013). Ecosystem respiration: Drivers of daily variability 676 and background respiration in lakes around the globe. *Limnology and Oceanography*, 677 58(3), 849-866. doi:10.4319/lo.2013.58.3.0849 678 Song, K., & Burgin, A. J. (2017). Perpetual Phosphorus Cycling: Eutrophication Amplifies 679 Biological Control on Internal Phosphorus Loading in Agricultural Reservoirs. 680 *Ecosystems*, 20(8), 1483-1493. doi:10.1007/s10021-017-0126-z 681 Staehr, P. A., Christensen, J. P. A., Batt, R. D., & Read, J. S. (2012). Ecosystem metabolism in a 682 stratified lake. Limnology and Oceanography, 57(5), 1317-1330. 683 doi:10.4319/lo.2012.57.5.1317 Stanley, E. H., Collins, S. M., Lottig, N. R., Oliver, S. K., Webster, K. E., Cheruvelil, K. S., & 684 685 Soranno, P. A. (2019). Biases in lake water quality sampling and implications for macroscale research. Limnology and Oceanography, 64(4), 1572-1585. 686 687 doi:10.1002/lno.11136 688 Stockwell, J. D., Doubek, J. P., Adrian, R., Anneville, O., Carey, C. C., Carvalho, L., ... Wilson, 689 H. L. (2020). Storm impacts on phytoplankton community dynamics in lakes. Global 690 Change Biology, 26(5), 2756-2784. doi:10.1111/gcb.15033 691 United States Environmental Protection Agency. (1993a). Determination of Nitrate-Nitrite by 692 Automated Colorimetry. Method 353.2 Revision 2.0. 693 United States Environmental Protection Agency. (1993b). Determination of Phosphorus by Semi-Automated Colorimetry. Method 365.1 Revision 2.0. 694 695 United States Environmental Protection Agency. (1993c). Determination of Total Kieldahl 696 Nitrogen by Semi-Automated Colorimetry Method 351.2, Revision 2.0. 697 Van de Bogert, M. C., Bade, D. L., Carpenter, S. R., Cole, J. J., Pace, M. L., Hanson, P. C., & 698 Langman, O. C. (2012). Spatial heterogeneity strongly affects estimates of ecosystem 699 metabolism in two north temperate lakes. Limnology and Oceanography, 57(6), 1689-700 1700. doi:10.4319/lo.2012.57.6.1689 701 Vilas, M. P., Marti, C. L., Adams, M. P., Oldham, C. E., & Hipsey, M. R. (2017). Invasive 702 Macrophytes Control the Spatial and Temporal Patterns of Temperature and Dissolved 703 Oxygen in a Shallow Lake: A Proposed Feedback Mechanism of Macrophyte Loss. 704 Frontiers in Plant Science, 8. doi:10.3389/fpls.2017.02097 705 Verpoorter, C., Kutser, T., Seekell, D. A., & Tranvik, L. J. (2014). A global inventory of lakes 706 based on high-resolution satellite imagery. Geophysical Research Letters, 41(18), 6396-707 6402. doi:10.1002/2014gl060641

708 Wilkinson, G. M., Carpenter, S. R., Cole, J. J., Pace, M. L., Batt, R. D., Buelo, C. D., & 709 Kurtzweil, J. T. (2018). Early warning signals precede cyanobacterial blooms in multiple 710 whole-lake experiments. Ecological Monographs, 88(2), 188-203. doi:10.1002/ecm.1286 711 Wu, X. D., Kong, F. X., Chen, Y. W., Qian, X., Zhang, L. J., Yu, Y., ... Xing, P. (2010). 712 Horizontal distribution and transport processes of bloom-forming Microcystis in a large 713 shallow lake (Taihu, China). Limnologica, 40(1), 8-15. doi:10.1016/j.limno.2009.02.001 714 Wynne, T. T., & Stumpf, R. P. (2015). Spatial and Temporal Patterns in the Seasonal 715 Distribution of Toxic Cyanobacteria in Western Lake Erie from 2002-2014. Toxins, 7(5), 716 1649-1663. doi:10.3390/toxins7051649 717 Zhou, Y. T., Obenour, D. R., Scavia, D., Johengen, T. H., & Michalak, A. M. (2013). Spatial and Temporal Trends in Lake Erie Hypoxia, 1987-2007. Environmental Science & 718 719 Technology, 47(2), 899-905. doi:10.1021/es303401b

721 Figures Captions

Figure 1. Sampling locations on a 65 m square grid of Swan Lake, a 40.5 hectare waterbody in western Iowa, USA. The main inlet to the lake and only outlet indicated with arrows. a) The location of the macrophyte beds of the two dominant species within the lake are shown on the map, with darker shading indicating the regions with the densest vegetation and the location of the high frequency sensor, b) the bathymetry of the lake and location of the lake in the state of Iowa, in reference to the United States of America.

728

729 **Figure 2.** The spatial pattern of each of the variables chlorophyll a (Chl, μ g L⁻¹), phycocyanin 730 (Phyco, µg L⁻¹), dissolved oxygen (DO, percent saturation), pH, and temperature (Temp, °C) for 731 each sampling event. The 98 sampling locations taken in a 65m grid (Figure 1) were interpolated 732 to a 25m grid using spatial inverse distance interpolation for visualization here. The color ramps 733 for each variable are scaled from the lowest to the highest value observed over the course of the 734 season across all sampling locations. The wind roses are the wind speeds (m s⁻¹; color ramp) and 735 direction the wind came from for the 24 hours prior to a sampling event. The concentric circles 736 are the frequency of winds from that direction for the 24 hour period (expressed as a percentage, 737 largest circle is 80% of the time). In the case of a longer "spoke", the greater amount of time the 738 wind was from that direction. The horizontal lines between DOY 170 and 177, and DOY 226 739 and 236 mark the two large precipitation events that occurred between those sampling dates.

740

Figure 3. Time series of the spatial coefficient of variation (CV) and spatial autocorrelation (AC;
local Moran's I) of the biologically-mediated variables in Swan Lake (same variable

abbreviations as Figure 2). The gray polygons indicate periods of algal bloom. The red line is the
time series of temperature local Moran's I for comparison.
Figure 4. Comparison of the mean (lines and points) and range (shaded polygon) of

- 747 measurements from the spatial sampling and fixed station measurements. The fixed station data
- 748 were trimmed to the period that spatial sampling occurred. A filled circle is used for the
- sampling dates when the means from the two sampling approaches were significantly different
- (p<0.05), and an open triangle is used for the sampling dates when the mean of the two
- approaches were not significantly different. The dark blue vertical lines indicate the dates of the
- two major precipitation events and the red dashed line in panel c is at 100% dissolved oxygen

saturation.

- 755 Figure 5. Standardized root mean squared errors (RMSE) of rarefaction analysis. Fit lines
- represent each sampling dates standardized RMSE (16 in total) and the gradient from light to
- 757 dark indicates first sampling event to last.