### 1 Capturing the spatial variability of algal bloom development in a shallow temperate lake

- 2 **Authors:** David Ortiz<sup>1,2\*</sup>, Grace Wilkinson<sup>1,2</sup>
- <sup>3</sup> <sup>1</sup>Department of Ecology, Evolution, and Organismal Biology, Iowa State University, Ames,
- 4 Iowa USA 50010
- 5 <sup>2</sup>Current Address: Center for Limnology, University of Wisconsin Madison, Madison,
- 6 Wisconsin USA 53706
- 7 \*Corresponding Author: dortiz4@wisc.edu
- 8
- 9 Keywords: spatial heterogeneity, spatial analysis, rarefaction analysis, macrophytes, eutrophic
- 10
- 11 This manuscript has been submitted for publication in Freshwater Biology. Please note that, despite having
- 12 undergone peer review, the manuscript has yet to be formally accepted for publication. Subsequent versions of this
- 13 manuscript may have slightly different content. If accepted, the final version of this manuscript will be available via
- 14 the 'Peer-reviewed Publication DOI' link on the right-hand side of this webpage. Please feel free to contact any of
- 15 the authors; we welcome feedback.

# 16 Abstract

17	1.	Algal blooms can have profound effects on the structure and function of aquatic
18		ecosystems and have the potential to interrupt valuable ecosystem services. Despite the
19		potential ecological and economic consequences of algal blooms, the spatial dynamics of
20		bloom development in spatially complex ecosystems such as shallow lakes remain poorly
21		characterized. Our goal was to evaluate the magnitude and drivers of spatial variability of
22		algal biomass, dissolved oxygen and pH over the course of a season, in a shallow lake in
23		order to better understand the spatial dynamics of algal blooms in these ecosystems.
24	2.	We sampled 98 locations in a small eutrophic lake on a 65m grid for several parameters
25		(chlorophyll a, phycocyanin, dissolved oxygen, pH, and temperature), weekly over 122
26		days. This was done to estimate the dynamics of variability and spatial autocorrelation
27		during the course of multiple bloom events. We also compared the spatial measurements
28		to a high frequency sensor deployed at a fixed station and estimated the optimal spatial
29		sampling resolution by performing a rarefaction analysis.
30	3.	Spatial heterogeneity of algal pigments was high, particularly during bloom events, and
31		this pattern and the overall severity of the bloom was not well captured with the fixed
32		station monitoring. The pattern of algal pigments and other limnologically important
33		variables (dissolved oxygen and pH) was related to the direction of prevailing winds 24
34		hours prior to sampling and likely enhanced by the shallow northern basin of the lake
35		where the main surface inlet was located. Additionally, a dense bed of floating-leaf
36		macrophytes contributed to local patchiness in all variables. Finally, from the rarefaction
37		analysis we found that minimal information about the mean state of the ecosystem was
38		gained after ~30 locations had been sampled.

4. This study revealed how spatially heterogeneous shallow lakes are over the course of a
single season, and that the magnitude of variability was highest during biologicallyintensive periods such as algal blooms. As such, continued research is needed across a
range of trophic conditions to better understand the structure of horizontal variability in
lakes. Overall, these data demonstrate the need for spatially-explicit monitoring to better
understand the dynamics and drivers of algal blooms in shallow lakes and to better
manage ecosystem services.

46 Introduction

47 Lakes are highly dynamic ecosystems that can undergo rapid physical and chemical 48 changes at an individual location, throughout their water column, and across the entire lake 49 surface at the scale of hours, days, seasons, and years (Laas et al., 2012; Read et al., 2011; Wynne & Stumpf, 2015). Quantifying heterogeneity in aquatic ecosystem structure and function 50 51 not only improves our understanding of lake ecology and the underlying mechanisms that drive 52 spatial and temporal heterogeneity, but also provides insights that improve management of these 53 ecosystems and the services they provide. With the development of sophisticated sensor 54 technology, high frequency measurements of variables such as dissolved oxygen and temperature 55 have helped limnologists grasp the scale of temporal heterogeneity in lakes (Carpenter et al., 56 2020; Chaffin et al., 2020; Cotterill et al., 2019). Detailed temporal monitoring has led to 57 advances in understanding several lake mechanisms such as diel cycles in primary production 58 (Solomon et al., 2013; Staehr et al., 2012), temperature effects on biogeochemical processes 59 (Medeiros et al., 2012), and early warnings of the transition to alternative stable states (Carpenter 60 et al., 2011; Wilkinson et al., 2018). Additionally, high frequency measurements have been used 61 to better understand heterogeneity over depth (vertical spatial heterogeneity) for important 62 processes such as stratification (Boehrer & Schultze, 2008; Read et al., 2011). Despite these 63 advances in understanding temporal and vertical heterogeneity, less is known about the dynamics 64 of horizontal spatial heterogeneity in the surface waters of lakes. 65 The vast majority of our understanding of lentic ecosystem structure and function comes

66 from single station sampling, with measurements taken through time over the deepest point in 67 the lake (Stanley et al., 2019). This location is usually selected to be representative of conditions 68 in the lake; however, the representativeness of a single location is likely to vary with regards to

69 the variable being measured and with time due to interacting forces such as wind, hydrology, 70 bathymetry, and biology (Chaffin et al., 2020; Schilder et al., 2013; Wu et al., 2010; Zhou., 71 2013). For example, ecosystem metabolism measured at dozens of locations for 10 days in two 72 north temperate lakes varied 1-2 orders of magnitude, with more than three-quarters of the 73 variability attributable to the measurement location within the lake (Van de Bogert et al., 2012). 74 Transect-based studies of reservoirs have revealed gradients in algae pigments, pH, and nutrients 75 with differences varying between 25%-180% within a waterbody (Moreno-Ostos et al., 2009; 76 Rychtecky & Znachor, 2011; Smith, 2018). Recently, satellite-based studies have demonstrated 77 the ability to detect spatial patterns at a high resolution for optical variables in large lakes (Lekki et al., 2019). Despite these advances, relatively few studies have quantified horizontal spatial 78 79 variability over time in lakes (Buttita et al. 2017, Vilas et al. 2017, Loken et al. 2019), hampering 80 our understanding of the magnitude of heterogeneity in variables important for managing water 81 quality and ecosystem services.

82 The development of algal blooms is expected to be a spatially heterogeneous 83 phenomenon (Buelo et al., 2018; Butitta, Carpenter et al., 2017; Serizawa et al., 2008) due to 84 both local heterogeneity in nutrient limitation, zooplankton grazing, and temperature (Davis et 85 al., 2009; Hansen et al., 1997) and population scale heterogeneity due to wind (George & 86 Heaney, 1978). Algal blooms can have a negative effect on ecosystem services, and therefore are 87 often a target for ecosystem monitoring and management. Some bloom-forming taxa, 88 particularly freshwater cyanobacteria, can produce toxins that rise to dangerous concentrations 89 for humans, pets, and livestock (Codd et al., 2005; Corbel et al., 2014). Additionally, the 90 mineralization of settling phytoplankton contributes to anoxic bottom waters, while intense 91 periods of primary production cause large variation in dissolved oxygen and pH (in poorly

buffered ecosystems) over the course of the day, which is stressful for aquatic organisms
(Gilbert, 2017; Landsberg, 2002). Furthermore, the perceived recreational value of lakes declines
when blooms form (Angradi et al., 2018), which in turn can negatively affect local economies
(Dodds et al., 2009). Despite the risk of economic loss, loss in biodiversity, and potential human
harm, the spatial dynamics of bloom development in spatially complex ecosystems such as
shallow lakes remain poorly characterized.

98 Shallow lakes have a large interface between the sediment and water relative to deeper 99 lakes, making them more susceptible to rapid changes in water residence time and nutrient inputs 100 (Christensen et al., 2013; Rennella & Quiros, 2006; Romo et al., 2013). Due to the expansive 101 littoral zones, shallow lakes can have large macrophyte beds which modify the light climate and 102 turbulence at the sediment-water interface (Andersen et al., 2017; Moller & Rordam, 1985; Vilas 103 et al., 2017). Many shallow lakes are also polymictic, experiencing multiple periods of 104 stratification followed by mixing during the ice-free season. During periods of water column 105 stability, some cyanobacteria taxa thrive, initiating blooms (Carey et al., 2012). Additionally, 106 episodic nutrient loading from the watershed during storm events (Carpenter et al., 2015; Kelly 107 et al., 2019), spatial gradients in nutrient availability due to stream inlets and morphology (e.g. 108 embayments), and wind-driven circulation (Schoen et al., 2014) can all contribute to spatial 109 heterogeneity of algal blooms over time in shallow lakes.

In order to better understand the spatial dynamics of algal blooms in shallow lakes, we performed intensive spatial sampling on Swan Lake (Iowa, USA), a spatially complex, shallow, hypereutrophic waterbody with a history of toxic cyanobacteria algal blooms. In addition to measuring algal pigments throughout the lake over the course of 122 days, we also measured temperature, dissolved oxygen, and pH. The spatial sampling captured two bloom events and

coincided with high frequency monitoring of the same variables using autonomous sensors
deployed at a fixed station (Ortiz et al. 2020). Using these data, we addressed the following
questions: 1) how does spatial variability of algae, dissolved oxygen, and pH change over the
course of a season, 2) are high frequency measurements at a fixed station an adequate
characterization of surface water dynamics in a shallow lake, and 3) what is the optimal spatial
sampling frequency to capture the mean state of a productive waterbody? Evaluating these
questions with data from a spatially complex, hypereutrophic lake will provide valuable
ecological and management-relevant insights into algal bloom dynamics.
Methods
Study Site
Swan Lake (42.0396, -94.8454) has an average depth of 2 m, surface area of 40.5
hectares, and a shoreline development index value of 1.54. The watershed is 350 hectares with
92% of the land in agricultural use. The estimated water residence time is approximately 1.5
years. During the ice-free period of 2018, Swan Lake had an average total phosphorus
years. During the ice-free period of 2018, Swan Lake had an average total phosphorus concentration of $280 \mu$ g L <sup>-1</sup> and a total nitrogen concentration of $1.61 m$ g L <sup>-1</sup> , making it
concentration of $280 \mu g L^{-1}$ and a total nitrogen concentration of 1.61 mg $L^{-1}$ , making it
concentration of $280 \mu g L^{-1}$ and a total nitrogen concentration of 1.61 mg $L^{-1}$ , making it hypereutrophic (Carlson, 1977). Total nitrogen was measured as the sum of total Kjeldahl
concentration of $280 \mu g L^{-1}$ and a total nitrogen concentration of 1.61 mg $L^{-1}$ , making it hypereutrophic (Carlson, 1977). Total nitrogen was measured as the sum of total Kjeldahl nitrogen (method 351.2 v2, US EPA, 1993c) and nitrate + nitrite measured using the cadmium

136 addition to seasonal algal blooms, Swan Lake also has non-continuous beds of American lotus

137 (Nelumbo lutea) and sago pondweed (Stuckenia pectinata) that peak in biomass in the latter half

of the summer and then begin senescing. The main surface inlet to the lake enters on the western
side and the outlet is at the southern edge of the waterbody (Figure 1). There are no known
springs feeding the lake.

141

142 Field Methods

143 The spatial sampling occurred approximately weekly from day of year (DOY) 142 to 144 DOY 264, encompassing the late spring, summer, and early autumn. A total of 16 spatial 145 sampling events occurred over the course of the 122 days. Measurements of chlorophyll a, 146 phycocyanin, temperature, dissolved oxygen saturation and pH were taken 0.25 m below the 147 surface at 98 sampling stations using a YSI Pro DSS multiparameter sonde (Yellow Springs 148 Instrument, Yellow Springs, OH) suspended over the side of a 3-meter long jon boat equipped 149 with an outboard motor. The sensors, which included the fluorometric Total Algae (chlorophyll a 150 and phycocyanin), optical dissolved oxygen, and Ag/AgCl pH sensors, were calibrated weekly 151 prior to each sampling event according to manufacturer instructions. The sampling stations were 152 laid out in a 65 x 65 m grid across the lake (Figure 1) with each location measured in the same 153 order (north to south) for each sampling event. This spatial resolution was selected to allow for 154 many sampling locations to be measured in a relatively short window of time, thereby 155 minimizing the chance that the differences observed between sampling locations was not due to 156 time of day. Measurements were taken between 10:00 and 14:00, except for the first two and last 157 three weeks when sampling lasted until 16:00. Beginning on DOY 177 when submerged 158 macrophytes could be identified from the jon boat, the presence or absence of submerged or 159 floating leaf macrophytes was noted at each sampling station during each sampling event. These 160 weekly presence/absence data were used to construct the macrophyte distributions in Figure 1.

161 The fixed station high frequency monitoring of Swan Lake was performed using a YSI 162 EXO2 (Yellow Springs Instrument, Yellow Springs, OH) multiparameter sonde equipped with 163 the same sensors as the YSI ProDSS used for the spatial sampling. The sonde recorded 164 measurements of chlorophyll a, phycocyanin, dissolved oxygen saturation, and pH every 15 minutes. The instrument was deployed on DOY 135 over the deepest point in the lake (3.8 m 165 166 deep), hanging approximately 0.5m below the surface, and removed on DOY 264 after the 167 spatial sampling event on that day. The fixed station sonde was monitored weekly for drift and 168 calibrated according to manufacturer instructions when indicated by the quality control algorithm 169 in the KorEXO software. Hourly precipitation, wind speed, and wind direction were collected at 170 the Arthur N. Neu Airport in Carroll, Iowa, located 4.5 km from the lake, as a part of the 171 National Oceanic and Atmospheric Automated Surface Observatory System. The meteorological 172 data were used to aid in the interpretation of spatial dynamics over the course of the summer. 173

174 Data analysis

175 Spatial heterogeneity can be quantified by calculating the spatial variance (e.g., 176 coefficient of variation; CV) or spatial autocorrelation (Moran's I, Moran, 1950). Increasing 177 spatial variance is indicative of increasing patchiness in the ecosystem, such as areas of high-178 density algal biomass and areas of low-density biomass within a lake. Spatial autocorrelation 179 accounts for the location of those patches within the ecosystem in relationship to each other. 180 Local Moran's I quantifies how similar the abundance of algae is at one location compared to the 181 density of surrounding neighbors. When measured over time for variables that are indices of 182 algal biomass (e.g., the pigments chlorophyll a and phycocyanin), both of these metrics of spatial 183 heterogeneity can provide insight into the dynamics of algal bloom development. In models of

algal blooms, both spatial variance and autocorrelation are expected to be high during the bloomperiod (Buelo et al., 2018).

186 Spatial autocorrelation (AC) and the coefficient of variation (CV) were calculated for 187 each variable on each sampling date in order to evaluate the dynamics of these parameters over 188 time. Prior to analysis, extreme outliers in the algal pigments were removed from the spatial 189 dataset as they were well outside the operating range of the Total Algae sensor or there was 190 known interference with the sensor resulting in an inaccurate measurement. This resulted in five 191 chlorophyll a and three phycocyanin measurements being removed out of 3,136 total pigment 192 measurements. The spatial CV is the standard deviation of all of the spatial measurements for a 193 variable on a given sampling date divided by the mean of those measurements, expressed as a 194 percent. Spatial AC was calculated as the average value of local Moran's I with a queen's 195 distance neighbor list (92 meters) with equal weight (1/n) on neighbors, as to not impose any 196 assumptions on possible spatial patterns in the variables. We limited our analysis to surrounding 197 neighbors because distances beyond this have not shown high spatial autocorrelation of algal 198 pigments under experimental conditions (Butitta et al. 2017). Local Moran's I values near 1.0 199 reflect high spatial AC within neighbors, zero indicates a random distribution, whereas spatial 200 AC values nearing -1.0 indicate a perfectly dispersed distribution (e.g. checkerboard pattern) in 201 the variable being measured. As the spatial variability in temperature is mediated by physical 202 processes, we used the dynamics and extent of the spatial AC of temperature as a benchmark to 203 visually compare the dynamics of spatial AC in the other biological variables. This allowed us to 204 tease apart the effect of physically-versus biologically-driven spatial patterns. Additionally, in 205 order to better visualize the spatial patterns in chlorophyll a, phycocyanin, temperature,

dissolved oxygen, and pH over the course of the season, the data were interpolated using inverse
distance weighting across a 25m grid (Figure 2).

208 In order to evaluate if high frequency measurements at a fixed station are an adequate 209 characterization of the surface water dynamics in a shallow lake, we compared the measurements 210 taken by the fixed station sonde during the same time period as a spatial sampling event. High 211 frequency data from the fixed station sonde was trimmed to the period that we sampled the lake 212 spatially. A t-test with a Bonferroni correction, to account for the multiple comparisons, was 213 performed to compare the distribution from the 98 sampling stations to the fixed station 214 measurements from the same day for each of the four biologically-mediated variables, 215 chlorophyll, phycocyanin, dissolved oxygen, and pH. In addition to comparing fixed sonde 216 values to the spatial sampling, we also used the spatial data to identify locations in the lake that 217 were consistently representative of mean conditions, and therefore ideal locations for fixed 218 station monitoring. We identified locations in the lake for each sampling event that had 219 measurements within the range of  $\pm$  one standard deviation from the mean for each biologically 220 mediated variable (all variables except temperature). We then collated these locations across all 221 sampling dates to identify which of the 98 sampling locations had measurements that most 222 consistently represented the mean conditions of the lake.

Finally, we performed a rarefaction analysis to evaluate the optimal spatial sampling frequency to capture the mean value of the biologically-mediated variables. This was done by randomly selecting *n* number of spatial sampling data points (n=2-97) during a sampling event, calculating the mean value from that subset, and then calculating the root mean square error (RMSE), comparing the mean estimate from the subset to the mean of all sampling points during that event. This calculation was repeated 1000 times for each value of *n*, and each iteration was

229	then averaged. The averaged RMSE values for each subset of $n$ were fit using a local polynomial
230	regression with a smoothing factor of 0.1 and each sampling event's RMSE curve was
231	standardized by subtracting the mean of all iterations ("global mean") from the mean at $n$
232	number of stations, to aid in visual comparison. The spatial data are available through (Ortiz &
233	Wilkinson, 2019) and the fixed station data are available in (Ortiz et al., 2019) and further
234	analyzed in Ortiz et al. (2020). All analyses were performed using R 4.0.2 (R Core Team, 2020)
235	using the gstat (Pebesma, 2004), rstatix (Kassambara, 2020), and sf packages (Pebesma, 2018).
236	
237	Results

238 There were two bloom events during the summer of 2018 in Swan Lake. The first bloom 239 occurred from DOY 156 – 184 and was dominated by the diatom Aulacoseira spp. based on a 240 sample taken on DOY 177 examined under a compound microscope at 400x magnification. The 241 phycocyanin concentrations on DOY 177 were the lowest during this first bloom period (Figure 242 2), and no cyanobacteria were identified in the sample. The second bloom, peaking on DOY 236, 243 was dominated by the cyanobacterium *Microcystis spp*. There were also two large precipitation events during the summer, occurring after sampling on DOY 170 and lasting through DOY 171, 244 245 and on DOY 232 (Figure 2; Supplemental Figure 1). The maximum wind speed recorded during 246 the first precipitation event was 10.8 m s<sup>-1</sup> coming from the southwest and 11.8 m s<sup>-1</sup> during the 247 second precipitation event coming from the southeast. During the first half of the summer (DOY 248 142 - 191) the prevailing winds 24 hours prior to the sampling events were from the south, 249 switched to being predominantly from the north from DOY 198 - 219, and then varied in 250 direction for the rest of the season (Figure 2). The median wind speed for the first period when 251 winds were out of the south was 3.6 m s<sup>-1</sup> (Figure 3b). When the winds switched to being

predominantly from the north between DOY 198 – 219, the median wind speed was lower at 2.5  $m s^{-1}$  (Figure S1).

254 Spatial dynamics

255 During the two bloom periods there was not a latitudinal or longitudinal trend in 256 chlorophyll a concentrations; instead, there were patches of high chlorophyll a concentration on 257 otherwise low-concentration dates (Figure 2). Unlike chlorophyll a, phycocyanin had a strong 258 latitudinal trend with higher concentrations in the northern portion (sample sites A1-G4) of the 259 lake during the first bloom. This spatial pattern is readily observed on DOY 184 but is also 260 noticeable for many of the sampling events during the first bloom (Figure 2). During sampling 261 events with a strong latitudinal gradient in phycocyanin (DOY 166 - 184, and 236) the mean 262 concentration in the northern portion of the lake was nearly double the concentration in the 263 southern portion of the lake (7.29 and  $3.76 \,\mu g \, L^{-1}$ , respectively). On these dates, the prevailing 264 winds 24 hours prior to the sampling event were out of the north (Supplemental Figure S1), yet 265 the lowest concentrations of phycocyanin were found in the southern portion of the lake. Even 266 when the lake was not blooming, there were patches of high concentrations of phycocyanin in 267 the northern portion of the lake (e.g., DOY 212), located among the densest patch of American 268 Lotus (Figure 1). The average phycocyanin concentrations at the sampling locations within the 269 American Lotus patch was higher than the average concentration in the rest of the lake for 14 of 270 the 16 sampling events (Figure S2).

The daytime saturation of dissolved oxygen varied the most out of the five variables monitored, ranging from borderline hypoxic (30% saturation) to supersaturated (up to 350%) (Figure 2). While the dissolved oxygen saturation increased near the peak of the bloom, the highest average saturation was on DOY 191, after the first bloom had collapsed. There was a

275 weak pattern over the course of the season of higher saturation in the northern portion of the 276 lake, similar to the distribution of higher phycocyanin concentrations. However, within the 277 northern portion of the lake, regions of low dissolved oxygen saturation formed in the surface 278 waters, particularly later in the summer (Figure 1). Beginning on DOY 198, the mean dissolved 279 oxygen concentration in the American lotus patch was consistently lower than the average for the 280 rest of the lake until DOY 250 (Figure S2). The distribution of pH also had a weak spatial pattern 281 during the summer, with slightly elevated values in the northern portion of the lake during the 282 first bloom (e.g. DOY 177; Figure 2). While pH was elevated at the onset of the first bloom 283 period from DOY 149 – 170, it was highest overall on DOY 191 and 198 after the first bloom 284 had collapsed. Unlike the other variables, temperature had a subtle south to north latitudinal 285 gradient with warmer temperatures in the southern portion of the lake and colder in the north 286 during the latter half of the summer (Figure 2). On average this difference between the northern 287 portion of the lake and the southern was 0.5°C. The warmest day of sampling was DOY 191. 288 Spatial variability in algal pigments during the first bloom event was low, with two 289 exceptions. There was an increase in the CV of chlorophyll a on the last day of the bloom (DOY 290 184; Figure 3a) that continued to increase as the bloom collapsed. There was also a temporary 291 increase in phycocyanin CV during the first bloom on DOY 177 (Figure 3b), coinciding with a 292 temporary decline in phycocyanin concentration across the lake. The CV of both algal pigments 293 was higher than the CV of temperature over the course of the entire sampling period. 294 Conversely, the CV of pH and dissolved oxygen were elevated during the first bloom 295 period, with pHCV declining and remaining low after the first bloom (Figure 3c) and dissolved 296 oxygen CV only temporarily declining after the first bloom (Figure 3d). Temperature had low 297 variability throughout the first bloom until DOY 177, when the lake began to heat up, peaking in

both temperature and spatial variability on DOY 191 (Figure 3e). Between the first and second
bloom, DOY 191-226, there was a decrease in spatial variability among the algal pigments and
pH as the bloom collapsed, while temperature and dissolved oxygen CV remained relatively high
and variable. During the second bloom period, CV was low for all variables except for
chlorophyll *a*. In general, the CV, of temperature and pH, expressed as a percentage, was an
order of magnitude lower than the other variables.

304 Spatial autocorrelation (AC), quantified as local Moran's I, did not fall substantially 305 below 0 for any of the variables and peaked at 0.79 among all variables (Figure 3). The highest 306 AC value for chlorophyll a and phycocyanin was during the first bloom event (Figure 3f, g); 307 however, phycocyanin AC also increased substantially during the second bloom. During the first 308 bloom, the AC of temperature varied similarly to both pigments' AC, particularly phycocyanin, 309 but became decoupled after the bloom collapsed. While the AC of temperature remained high 310 during the inter-bloom period, the AC of the pigments was substantially lower. Conversely, the 311 dynamics of AC of temperature, dissolved oxygen and pH remained coupled throughout the 312 summer (Figure 3h, i). Dissolved oxygen saturation and pH both increased in AC during the first 313 bloom and then declined throughout the rest of the season with the exception of a minor increase 314 in AC during the second bloom event.

315

#### 316 Fixed station versus spatial sampling

317 There were a greater number of days with a significant difference between the spatial and 318 fixed station measurements than days in which the data sets were not significantly different 319 (Figure 4). Among all 64 comparisons (4 variables × 16 sampling events), the spatial and fixed 320 station data sets had a means that were not significantly different 37.5% of the time.

321 Phycocyanin had the greatest number of events with similar values, with 7 of the 16 sampling 322 events having non-statistically different mean values measured spatially and at the fixed station 323 (Figure 4b). These occurrences were mainly during non-bloom periods. However, even when the 324 mean phycocyanin values were similar between the sampling methods on a given day, the range 325 of values captured by the fixed station was five times less than the variability captured in the 326 spatial data. This pattern of infrequent occurrences of similar mean values between the two 327 methods during non-bloom periods and a diminished range in the fixed station data, was shared 328 to a degree, among the other three variables as well. Interestingly, dissolved oxygen saturation 329 only had 5 out of the 16 events with means that were not significantly different, all of which 330 occurred when the lake was above 100% saturation (Figure 4c).

While a majority of the comparisons between the fixed station and spatial data indicate that the algal pigments had a larger range of values in the spatial data, there were a handful of instances where the opposite was true. During the first bloom, the fixed station sonde measured a wide range of chlorophyll *a* concentrations and had a higher mean chlorophyll *a* for all dates (Figure 4a). Similarly, we observed higher mean phycocyanin at the fixed station sonde on DOY 156, 166, 177, 191, and 219 (Figure 4b). However, this pattern did not hold true for dissolved oxygen or pH (Figure 5c, d).

The spatial sampling sites that most consistently captured the mean values in the lake on a given sampling date were in the northwest portion of the lake, near the inlet. The best performing site for all variables was site E3, just west of the American lotus patch and adjacent to a bed of sago pondweed (Figure 1). The four biologically-mediated variables from sample site E3 were within the mean (± standard deviation) range of all of the spatial measurements 95% of the time. The second best performing location was in the middle of the American lotus patch, site

D4, with the values from this site being within the mean (± standard deviation) range 92% of the
time. The site where the fixed station was located, site H2, was only within the mean (± standard
deviation) range 58% of the time.

347

348 Optimal Spatial Resolution

349 In order to evaluate the spatial sampling resolution needed to capture the mean state of 350 the surface water on a given day, we performed a rarefaction analysis for each variable and each 351 sampling event, calculating the root mean squared error (RMSE) of a subset of sampling 352 locations compared to the mean value of all 98 measurements that day. The plateaus of the 353 RMSE curves from the rarefaction analysis were used to evaluate the smallest number of spatial 354 sampling locations needed to capture the mean across the lake during that sampling event (Figure 355 5). Additionally, we also evaluated the temporal pattern of the minimum number of sampling 356 locations needed to capture the mean.

357 Mean values were underestimated for all variables on all sampling dates when there were 358 less than 10 sampling stations (Figure 5). However, the severity of the underestimation differed 359 among the variables. The rarefaction analysis for chlorophyll a indicated that 10-30 sampling 360 locations was sufficient for capturing the mean chlorophyll a in Swan Lake, otherwise the mean 361 concentration would be under estimated (Figure 5a). When an algal bloom was occurring it took 362 more sampling locations to near the mean chlorophyll a concentration on that date. However, 363 when the bloom was particularly patchy during development (DOY 226) or collapse (DOY 191), 364 including a larger number of sampling locations led to overestimating the mean chlorophyll a 365 concentration as locations with high concentrations were over-represented in the data set. There 366 were similar patterns in phycocyanin RMSE with most sampling dates plateauing between 20 -

367 30 sampling locations with a few exceptions (Figure 5b). For DOYs 156-170 (rise of the first 368 bloom) and 212, at least 60 sampling locations were needed to capture the overall mean in 369 phycocyanin for that sampling date. Dissolved oxygen saturation and pH were generally well 370 characterized by approximately 10-15 sampling locations as both had a majority of dates in 371 which the RMSE curves plateaued at that spatial sampling resolution (Figure 5c, d). However, at 372 the beginning (DOY 154), peak (DOY 184), and end (DOY 205) of the first bloom, twice as 373 many sampling locations were needed to capture the mean dissolved oxygen. Only two dates 374 required more sampling locations for pH to capture the mean, DOY 177 and 198, which 375 plateaued at approximately 40 sampling locations. The largest RMSE were observed during 376 bloom conditions for all variables: DOY 177 had the largest error for phycocyanin and pH, while 377 the largest RMSE was on DOY 184 for dissolved oxygen and on DOY 236 for chlorophyll a 378 (Figure 5).

379

### 380 **Discussion**

381 The spatial heterogeneity of water quality parameters was highly dynamic in Swan Lake, 382 a shallow, hypereutrophic, temperate waterbody. The temporal dynamics in heterogeneity were 383 driven in part by the two blooms, the peaks of which were preceded by large precipitation events. 384 These rain events could have delivered nutrients from the agriculturally dominated watershed 385 into the lake from the northern inlet, helping to fuel the subsequent algal blooms and the spatial 386 patterns observed during blooms (Stockwell et al. 2020). However, there are also a number of 387 other factors that likely contributed to the spatial variability and pattern during and following 388 these bloom events, including the prevailing wind direction prior to sampling, the bathymetry of

the basin and location of the surface inlet, and the potential for dense macrophyte beds tocontribute to local patchiness.

391 The spatial patterns that the algal blooms created were consistent with the expectations 392 from previous modeling and experimental work that spatial AC increases as algal blooms 393 develop (Buelo et al., 2018; Butitta et al., 2017; Serizawa et al., 2008). This pattern was the 394 strongest for phycocyanin, evident by the strong latitudinal gradient in concentrations during the 395 bloom periods. The sampling dates with phycocyanin concentration gradients (e.g., DOY 166, 396 177, 184, 236) coincided with persistent winds from the south 24 hours prior to the sampling 397 event, which likely resulted in the higher concentration of algal cells in the northern portion of 398 the lake. The effect of persistent wind directions influencing the distribution of cyanobacteria has 399 also been documented in other shallow eutrophic lakes (Wu et al. 2010). The shallow sediments 400 of the northern basin were also likely a source of akinete recruitment (Karlsson-Elfgren and 401 Brunburg, 2004), further contributing to the higher concentrations of phycocyanin in the northern 402 portion of the lake during the first bloom. Augmented nutrient availability in the northern part of 403 the lake due to external loading from the watershed through the surface inlet and internal loading 404 from the sediments overlain by an unstratified water column (Song and Burgin 2017) may have 405 further amplified the phytoplankton gradient, particularly following precipitation events. Finally, 406 the tendency of the dominant cyanobacteria taxa *Microcystis spp.* to form surface scums likely 407 enhanced the spatial patterns observed with our surface sampling approach.

The sampling dates with a strong gradient of phytoplankton concentrations from north to south also resulted in north-south gradients in water chemistry. On these dates, both dissolved oxygen and pH formed a gradient of high values in the northern portion of the lake and lower values in the south, which would be expected with greater primary production where

412 phytoplankton concentrations were highest. The spatial patterns in the surface water chemistry 413 demonstrate how phytoplankton spatial distribution, driven by wind, can create hot spots and 414 moments of biogeochemical activity within lakes (McClain et al. 2003) that may be missed with 415 traditional, single-station sampling. The dense patch of floating leaf American lotus macrophytes 416 also created a hot spot of biogeochemical activity.

417 Macrophyte beds can have a large local influence on water chemistry by inducing 418 stratification, decreasing flow and trapping particles, and modifying the light environment 419 (Green, 2006; Vilas et al., 2017). For 87.5% of the season the phycocyanin concentrations were 420 higher in the bed of American lotus than concentrations elsewhere in the lake. In fact, even on 421 sampling dates when phycocyanin concentrations were otherwise low (e.g., DOY 212), the 422 American lotus patch can be identified based on the phycocyanin concentrations that are nearly 423 twice as high as the rest of the lake. We hypothesize that the macrophyte patch allowed for 424 microstratification in the water column and reduced wind-driven flow. These physical conditions 425 are likely to favor cyanobacteria dominance and the formation of surface scums. Similarly, the 426 dissolved oxygen concentrations in the American lotus patch became consistently lower than the 427 rest of the lake later in the summer, likely due to the plants beginning to senesce, creating a hot 428 spot of decomposition, decreasing both dissolved oxygen and pH (Vilas et al., 2017). While there 429 isn't strong evidence in the data that the other submerged macrophyte beds had a similarly strong 430 effect on water chemistry, the data from the American lotus patch illustrates how macrophytes 431 can contribute to local patchiness and overall spatial heterogeneity.

432

433 Considerations for Monitoring

434 The variables that we measured in this study are often the target of water quality 435 monitoring as the dynamics of these variables coincide with changes in ecosystem function and 436 services. Monitoring is often performed at a fixed station over time to capture the dynamics of 437 the ecosystem, but this strategy could potentially result in missed information about the 438 ecosystem's behavior. While the temporal dynamics of all the variables were synchronous 439 between the fixed station and spatial sampling data sets in Swan Lake, our conclusions regarding 440 the magnitude of the blooms and variability in the lake's structure would have been substantially 441 different relying solely on the fixed station data. Among the four biologically-mediated 442 variables, only 37.5% of the fixed station estimates of the mean state of the lake statistically 443 matched the estimate from the spatial sampling. The vast majority of those instances (96%) 444 occurred during non-bloom periods, which also coincided with lower wind speed conditions, no 445 prevailing wind direction, and no major precipitation events. The large difference between the 446 spatial sampling and fixed station measurements of algal pigments during blooms was likely 447 driven, in part, by the depth of the sensors at the fixed station and the variable accumulation of 448 cyanobacteria at the surface of the lake dependent upon environmental conditions and the 449 dominant taxa (Chaffin et al., 2020). It is clear from our data that during periods of heightened 450 biological activity such as blooms, fixed station monitoring is unlikely to be representative of the 451 mean ecosystem state in shallow lakes.

Despite the high degree of horizontal spatial variability that has been documented in this study and others (Loken et al. 2019, Van de Bogert et al. 2012, Buttita et al. 2017), fixed station designs are widely used in water quality monitoring programs. In Swan Lake, we determined that the historical location for water quality monitoring, where the fixed station sensors were deployed, was one of the least-representative locations for mean conditions in the lake. Given the

457 hypereutrophic state of the lake, the most immediate management concerns are toxic 458 cyanobacteria blooms and summer fish kills due to low dissolved oxygen. Yet, the mean value of 459 these variables (phycocyanin and dissolved oxygen) across the lake were only captured by the 460 fixed station sensors 58% of the time. While selecting a fixed station site for high frequency 461 sensor deployment includes many considerations including the location of previous data 462 collection and management needs, based on our analysis we would advise performing a spatial 463 survey to identify if and when the fixed station site is representative of mean conditions in the 464 lake. A complementary spatial survey will help contextualize the fixed station dynamics and 465 provide additional, management-relevant information about the lake. 466 It's also important to consider the trade-offs between high frequency fixed station 467 monitoring and higher resolution, but less frequent spatial monitoring. High frequency 468 monitoring at a single station provides insight into ecosystem function such as metabolism 469 (Staehr et al., 2012), early warnings of impending regime shifts (Carpenter et al., 2011; 470 Wilkinson et al., 2018), and crucial information on diel variability in limnological conditions 471 (Andersen et al., 2017). However, as we observed in Swan Lake, the spatial variability within a 472 given day often exceeds the temporal variability at a single point in a shallow lake. Without the 473 spatial sampling snapshots, we would have underestimated the magnitude of the algal blooms, 474 hampering our limnological understanding of the ecosystem's functioning and impeding our 475 ability to accurately estimate rates such as methane emissions on a global scale (DelSontro et al. 476 2018). 477 From a practical stand point, the understanding gleaned from the spatial sampling could

477 From a practical stand point, the understanding gleaned from the spatial sampling could
478 help managers design targeted algal toxin monitoring or management interventions to help
479 control fish habitat quality in persistently hypoxic areas (Bardshaw et al., 2015). However, the

480 time and cost investment in repeated spatial sampling at the resolution performed in this study 481 may not be feasible for both research and management programs. The rarefaction analysis we 482 performed for all four of the key water quality monitoring variables revealed that minimal 483 information was gained after ~30 locations were sampled across many conditions and variables. 484 Often 12-20 sample locations across the 40.5 ha lake (or a 1-2 samples per hectare) was 485 sufficient to capture the spatial variability within the lake, with a few exceptions. These 486 exceptions occurred during times of higher variability such as when the blooms were just starting 487 or when the bloom began to collapse. The need for a higher spatial resolution during bloom 488 events to fully capture their variability has also been found using remote sensing techniques in 489 other, larger lakes (Lekki et al., 2019). As the spatial resolution of remote sensing technologies 490 continues to improve, it may become more cost effective to capture the spatial heterogeneity of 491 algal pigments in small lakes over time. However, one of the benefits of manual spatial sampling 492 is being able to pair other measurements such as dissolved oxygen, pH, and nutrients (e.g., 493 nitrate; Loken et al., 2018; Pellerin et al., 2016) with information on the distribution of algal 494 biomass.

495 Our intensive spatial monitoring of a shallow, hypereutrophic lake revealed how spatially 496 heterogeneous shallow lakes are over the course of a single season and allowed us to tease apart 497 the drivers of that spatial heterogeneity. We found that variability was greatest during 498 biologically-intensive periods, such as during algal blooms and in dense floating-leaf macrophyte 499 beds, and that failure to capture this variability would have hampered our understanding of the 500 ecosystem's functioning and overall mean state. Small lakes such as Swan Lake dominate the 501 global distribution of waterbodies (Verpoorter et al. 2014). Adequately capturing and 502 characterizing the magnitude of variability in production of these waterbodies is important given

503	their role in mediating global nutrient cycles (Downing et al. 2010, Biddanda et al. 2017),
504	especially methane emissions (DelSontro et al. 2018, Loken et al. 2019). Our data provided an
505	estimate of the spatial resolution needed to capture the dynamics in ecosystems similar to Swan
506	Lake and a method which could be readily adapted to other ecosystems. While our results
507	provide new understanding of the magnitude and temporal dynamics of spatial heterogeneity in
508	shallow lakes, continued investigation of horizontal spatial heterogeneity in a range of aquatic
509	ecosystems, from oligotrophic to eutrophic, is needed to better understand the structure and
510	drivers of horizontal spatial variability in lakes.
511	
512	Acknowledgments
513	We would like to thank Ryan Wagner, Ellen Albright, Rachel Fleck and Tyler Butts for
514	assistance with data collection and instrument deployment and collection. We would also like to
515	thank two anonymous reviewers and the editor for constructive comments that improved the
516	manuscript. Funding was provided by the Center for Global and Regional Environmental
517	Research, the Iowa State University Graduate Minority Assistantship Program, and the Iowa
518	State University Graduate Research Assistantship Match Program.
519	
520	References
521 522 523 524	American Public Health Association (APHA), American Water Works Association (AWWA), and the Water Environmental Federation (WEF). 1998. Standard Methods for Examinations of Water and Wastewater, 20th ed. United Book Press, Inc. Baltimore, Maryland.
525 526 527 528 529 530	<ul> <li>Andersen, M. R., Kragh, T., &amp; Sand-Jensen, K. (2017). Extreme diel dissolved oxygen and carbon cycles in shallow vegetated lakes. <i>Proceedings of the Royal Society B-Biological Sciences</i>, 284(1862), doi:10.1098/rspb.2017.1427</li> <li>Angradi, T. R., Ringold, P. L., &amp; Hall, K. (2018). Water clarity measures as indicators of recreational benefits provided by US lakes: Swimming and aesthetics. <i>Ecological Indicators</i>, 93, 1005-1019. doi:10.1016/j.ecolind.2018.06.001</li> </ul>

- Bardshaw, E. L., Allen, M. S., & Netherland, M. (2015). Spatial and temporal occurrence of
   hypoxia influences fish habitat quality in dense Hydrilla verticillata. *Journal of Freshwater Ecology*, 30(4), 491-502.
- Biddanda, B. A. (2017). Global significance of the changing freshwater carbon cycle. *EOS*, *98*,
   doi:10.1029/2017EO069751
- Boehrer, B., & Schultze, M. (2008). Stratification of lakes. *Reviews of Geophysics*, 46(2),
   doi:10.1029/2006rg000210
- Buelo, C. D., Carpenter, S. R., & Pace, M. L. (2018). A modeling analysis of spatial statistical
  indicators of thresholds for algal blooms. *Limnology and Oceanography Letters*, 3(5),
  384-392. doi:10.1002/lol2.10091
- Butitta, V. L., Carpenter, S. R., Loken, L. C., Pace, M. L., & Stanley, E. H. (2017). Spatial early
  warning signals in a lake manipulation. *Ecosphere*, 8(10), doi:10.1002/ecs2.1941
- 543 Carlson, R. E. (1977). Trophic State Index for Lakes. *Limnology and Oceanography*, 22(2), 361 544 369. doi:10.4319/lo.1977.22.2.0361
- Carpenter, S., Booth, E., Kucharik, C., & Lathrop, R. (2015). Extreme daily loads: role in annual
  phosphorus input to a north temperate lake. *Aquatic Sciences*, 77(1), 71-79.
  doi:10.1007/s00027-014-0364-5
- 548 Carpenter, S. R., Arani, B. M. S., Hanson, P. C., Scheffer, M., Stanley, E. H., & Van Nes, E.
  549 (2020). Stochastic dynamics of Cyanobacteria in long-term high-frequency observations
  550 of a eutrophic lake. *Limnology and Oceanography Letters*, 5(5), 331-336.
  551 doi:10.1002/lol2.10152
- Carpenter, S. R., Cole, J. J., Pace, M. L., Batt, R., Brock, W. A., Cline, T., . . . Weidel, B. (2011).
   Early Warnings of Regime Shifts: A Whole-Ecosystem Experiment. *Science*, *332*(6033),
   1079-1082. doi:10.1126/science.1203672
- Chaffin, J. D., Kane, D. D., & Johnson, A. (2020). Effectiveness of a fixed-depth sensor
   deployed from a buoy to estimate water-column cyanobacterial biomass depends on wind
   speed. *Journal of Environmental Sciences*, 93, 23-29, doi:10.1016/j.jes.2020.03.003
- Christensen, J. P. A., Sand-Jensen, K., & Staehr, P. A. (2013). Fluctuating water levels control
   water chemistry and metabolism of a charophyte-dominated pond. *Freshwater Biology*,
   58(7), 1353-1365. doi:10.1111/fwb.12132
- 561 Codd, G. A., Morrison, L. F., & Metcalf, J. S. (2005). Cyanobacterial toxins: risk management
  562 for health protection. *Toxicology and Applied Pharmacology*, 203(3), 264-272.
  563 doi:10.1016/j.taap.2004.02.016
- Corbel, S., Mougin, C., & Bouaicha, N. (2014). Cyanobacterial toxins: Modes of actions, fate in
   aquatic and soil ecosystems, phytotoxicity and bioaccumulation in agricultural crops.
   *Chemosphere*, 96, 1-15. doi:10.1016/j.chemosphere.2013.07.056
- 567 Cotterill, V., Hamilton, D. P., Puddick, J., Suren, A., & Wood, S. A. (2019). Phycocyanin
  568 sensors as an early warning system for cyanobacteria blooms concentrations: a case study
  569 in the Rotorua lakes. *New Zealand Journal of Marine and Freshwater Research*, 53(4),
  570 555-570. doi:10.1080/00288330.2019.1617322
- Davis, T. W., Berry, D. L., Boyer, G. L., & Gobler, C. J. (2009). The effects of temperature and
  nutrients on the growth and dynamics of toxic and non-toxic strains of Microcystis during
  cyanobacteria blooms. *Harmful Algae*, 8(5), 715-725. doi:10.1016/j.hal.2009.02.004
- 574 DelSontro, T., Beaulieu, J. J., & Downing, J. A. (2018). Greenhouse gas emissions from lakes
   575 and impoundments: Upscaling in the face of global change. *Limnology and* 576 Oceanography Letters, 3(3), 64-75. doi:10.1002/lol2.10073

577 Dodds, W. K., Bouska, W. W., Eitzmann, J. L., Pilger, T. J., Pitts, K. L., Riley, A. J., . . . 578 Thornbrugh, D. J. (2009). Eutrophication of US Freshwaters: Analysis of Potential 579 Economic Damages. Environmental Science & Technology, 43(1), 12-19. 580 doi:10.1021/es801217g 581 Downing, J. A. 2010. Emerging global role of small lakes and ponds: little things mean a lot. 582 *Limnetica* 29(1), 9-24. 583 George, D. G., & Heaney, S. I. (1978). Factors influencing spatial-distribution of phytoplankton 584 in a small productive lake. Journal of Ecology, 66(1), 133-155. doi:10.2307/2259185 585 Gilbert, P. M. (2017). Eutrophication, harmful algae and biodiversity - Challenging paradigms in 586 a world of complex nutrient changes. Marine Pollution Bulletin, 124(2), 591-606. 587 doi:10.1016/j.marpolbul.2017.04.027 Green, J. C. (2006). Effect of macrophyte spatial variability on channel resistance. Advances in 588 589 Water Resources, 29(3), 426-438. doi:10.1016/j.advwatres.2005.05.010 590 Hansen, A. M., Andersen, F. O., & Jensen, H. S. (1997). Seasonal pattern in nutrient limitation 591 and grazing control of the phytoplankton community in a non-stratified lake. Freshwater 592 *Biology*, *37*(3), 523-534. doi:10.1046/j.1365-2427.1997.00182.x 593 Kassambara, A. (2020). rstatix: Pipe-Friendly Framework for Basic Statistical Tests. R package 594 version 0.6.0. Retrieved from https://cran.r-project.org/package=rstatix 595 Karlsson-Elfgren, I., & Brunberg, A. K. (2004). The importance of shallow sediments in the 596 recruitment of Anabaena and Aphanizomenon (Cyanophyceae). Journal of Phycology, 597 40(5), 831-836. doi:10.1111/j.1529-8817.2004.04070.x 598 Kelly, P. T., Renwick, W. H., Knoll, L., & Vanni, M. J. (2019). Stream Nitrogen and Phosphorus 599 Loads Are Differentially Affected by Storm Events and the Difference May Be 600 Exacerbated by Conservation Tillage. Environmental Science & Technology, 53(10), 601 5613-5621. doi:10.1021/acs.est.8b05152 602 Laas, A., Noges, P., Koiv, T., & Noges, T. (2012). High-frequency metabolism study in a large 603 and shallow temperate lake reveals seasonal switching between net autotrophy and net 604 heterotrophy. Hydrobiologia, 694(1), 57-74. doi:10.1007/s10750-012-1131-z 605 Landsberg, J. H. (2002). The effects of harmful algal blooms on aquatic organisms. Reviews in 606 Fisheries Science, 10(2), 113-390. doi:10.1080/20026491051695 607 Lekki, J., Deutsch, E., Sayers, M., Bosse, K., Anderson, R., Tokars, R., & Sawtell, R. (2019). 608 Determining remote sensing spatial resolution requirements for the monitoring of harmful 609 algal blooms in the Great Lakes. Journal of Great Lakes Research, 45(3), 434-443. 610 doi:10.1016/j.jglr.2019.03.014 611 Loken, L. C., Crawford, J. T., Dornblaser, M. M., Striegl, R. G., Houser, J. N., Turner, P. A., & 612 Stanley, E. H. (2018). Limited nitrate retention capacity in the Upper Mississippi River. 613 Environmental Research Letters, 13(7). doi:10.1088/1748-9326/aacd51 614 Loken, L. C., Crawford, J. T., Schramm, P. J., Stadler, P., Desai, A. R. & Stanley, E. H. (2019). 615 Large Spatial and Temporal Variability of Carbon Dioxide and Methane in a Eutrophic 616 Lake. Journal of Geophysical Research-Biogeosciences, 124(7), 2248-2266. 617 10.1029/2019jg005186 618 McClain, M. E., Boyer, E. W., Dent, C. L., Gergel, S. E., Grimm, N. B., Groffman, P. M., ... 619 Pinay, G. (2003). Biogeochemical hot spots and hot moments at the interface of terrestrial and aquatic ecosystems. *Ecosystems*, 6(4), 301-312. doi:10.1007/s10021-003-0161-9 620 Medeiros, A. S., Biastoch, R. G., Luszczek, C. E., Wang, X. A., Muir, D. C. G., & Quinlan, R. 621 622 (2012). Patterns in the limnology of lakes and ponds across multiple local and regional

623 environmental gradients in the eastern Canadian Arctic. Inland Waters, 2(2), 59-76. 624 doi:10.5268/iw-2.2.427 625 Moller, T. R., & Rordam, C. P. (1985). Species numbers of vascular plants in relation to area, 626 isolation and age of ponds in Denmark. Oikos, 45(1), 8-16. doi:10.2307/3565216 627 Moran, P. A. P. (1950). Notes on Continuous Stochastic Phenomena. Biometrika, 37(1-2), 17-23. 628 doi:10.2307/2332142 629 Moreno-Ostos, E., Cruz-Pizarro, L., Basanta, A., & George, D. G. (2009). Spatial Heterogeneity 630 of Cyanobacteria and Diatoms in a Thermally Stratified Canyon-Shaped Reservoir. International Review of Hydrobiology, 94(3), 245-257. doi:10.1002/iroh.200811123 631 632 Ortiz, D., Palmer, J., & Wilkinson, G. (2019). Hypereutrophic lake sensor data during summer 633 algae blooms in Iowa, USA, 2014 - 2018 ver 1. Environmental Data Initiative. 634 doi:10.6073/pasta/30070d41fbcdf36387f33d9108f570f8 Ortiz, D., & Wilkinson, G. (2019). Hypereutrophic lake spatial sensor data during summer 635 636 bloom, Swan Lake, Iowa, USA 2018 ver 1. Environmental Data Initiative. 637 doi:10.6073/pasta/2c0ca177438a3d422925811514e86cd8 Ortiz, D., Palmer, J., & Wilkinson, G. (2020). Detecting changes in statistical indicators of 638 639 resilience prior to algal blooms in shallow eutrophic lakes. *Ecosphere*, 11(10). 640 doi:10.1002/ecs2.3200 641 Pace, M. L., Batt, R. D., Buelo, C. D., Carpenter, S. R., Cole, J. J., Kurtzweil, J. T., & Wilkinson, 642 G. M. (2017). Reversal of a cyanobacterial bloom in response to early warnings. 643 Proceedings of the National Academy of Sciences of the United States of America, 644 114(2), 352-357. doi:10.1073/pnas.1612424114 645 Pebesma, E. J. (2018). Simple Features for R: Standardized Support for Spatial Vector Data. R 646 Journal, 10(1), 439-446. 647 Pebesma, E. J. (2004). Multivariable geostatistics in S: the gstat package. Computers & 648 Geosciences, 30(7), 683-691. doi:10.1016/j.cageo.2004.03.012 649 Pellerin, B. A., Stauffer, B. A., Young, D. A., Sullivan, D. J., Bricker, S. B., Walbridge, M. R., 650 ... Shaw, D. M. (2016). Emerging tools for continuous nutrient monitoring networks: 651 sensors advancing science and water resources protection. Journal of the American Water 652 Resources Association, 52(4), 993-1008. doi:10.1111/1752-1688.12386 653 R Core Team. (2020). R: A language and environment for statistical computing. R Foundation 654 for Statistical Computing, Vienna, Austria. Retrieved from https://www.R-project.org/ 655 Read, J. S., Hamilton, D. P., Jones, I. D., Muraoka, K., Winslow, L. A., Kroiss, R., ... Gaiser, E. 656 (2011). Derivation of lake mixing and stratification indices from high-resolution lake 657 buoy data. Environmental Modelling & Software, 26(11), 1325-1336. 658 doi:10.1016/i.envsoft.2011.05.006 659 Rennella, A. M., & Quiros, R. (2006). The effects of hydrology on plankton biomass in shallow 660 lakes of the Pampa Plain. Hydrobiologia, 556, 181-191. doi:10.1007/s10750-005-0318-y 661 Romo, S., Soria, J., Fernandez, F., Ouahid, Y., & Baron-Sola, A. (2013). Water residence time 662 and the dynamics of toxic cyanobacteria. Freshwater Biology, 58(3), 513-522. 663 doi:10.1111/j.1365-2427.2012.02734.x 664 Rychtecky, P., & Znachor, P. (2011). Spatial heterogeneity and seasonal succession of 665 phytoplankton along the longitudinal gradient in a eutrophic reservoir. Hydrobiologia, 663(1), 175-186. doi:10.1007/s10750-010-0571-6 666 667 Schilder, J., Bastviken, D., van Hardenbroek, M., Kankaala, P., Rinta, P., Stotter, T., & Heiri, O. 668 (2013). Spatial heterogeneity and lake morphology affect diffusive greenhouse gas

669 emission estimates of lakes. Geophysical Research Letters, 40(21), 5752-5756. 670 doi:10.1002/2013gl057669 671 Schoen, J. H., Stretch, D. D., & Tirok, K. (2014). Wind-driven circulation patterns in a shallow 672 estuarine lake: St Lucia, South Africa. Estuarine Coastal and Shelf Science, 146, 49-59. doi:10.1016/j.ecss.2014.05.007 673 674 Serizawa, H., Amemiya, T., & Itoh, K. (2008). Patchiness in a minimal nutrient - phytoplankton 675 model. Journal of Biosciences, 33(3), 391-403. doi:10.1007/s12038-008-0059-y 676 Smith, C., D. (2018). Temporal and Spatial Monitoring of Cyanobacterial Blooms at Willow 677 Creek Reservoir, North-Central Oregon. Retrieved from U.S. Geological Survey 678 Scientific Investigations Report: 679 Solomon, C. T., Bruesewitz, D. A., Richardson, D. C., Rose, K. C., Van de Bogert, M. C., 680 Hanson, P. C., ... Zhu, G. W. (2013). Ecosystem respiration: Drivers of daily variability 681 and background respiration in lakes around the globe. *Limnology and Oceanography*, 682 58(3), 849-866. doi:10.4319/lo.2013.58.3.0849 683 Song, K., & Burgin, A. J. (2017). Perpetual Phosphorus Cycling: Eutrophication Amplifies 684 Biological Control on Internal Phosphorus Loading in Agricultural Reservoirs. 685 *Ecosystems*, 20(8), 1483-1493. doi:10.1007/s10021-017-0126-z 686 Staehr, P. A., Christensen, J. P. A., Batt, R. D., & Read, J. S. (2012). Ecosystem metabolism in a 687 stratified lake. Limnology and Oceanography, 57(5), 1317-1330. 688 doi:10.4319/lo.2012.57.5.1317 Stanley, E. H., Collins, S. M., Lottig, N. R., Oliver, S. K., Webster, K. E., Cheruvelil, K. S., & 689 690 Soranno, P. A. (2019). Biases in lake water quality sampling and implications for 691 macroscale research. Limnology and Oceanography, 64(4), 1572-1585. 692 doi:10.1002/lno.11136 693 Stockwell, J. D., Doubek, J. P., Adrian, R., Anneville, O., Carey, C. C., Carvalho, L., ... Wilson, 694 H. L. (2020). Storm impacts on phytoplankton community dynamics in lakes. Global 695 Change Biology, 26(5), 2756-2784. doi:10.1111/gcb.15033 United States Environmental Protection Agency. (1993a). Determination of Nitrate-Nitrite by 696 697 Automated Colorimetry. Method 353.2 Revision 2.0. 698 United States Environmental Protection Agency. (1993b). Determination of Phosphorus by 699 Semi-Automated Colorimetry. Method 365.1 Revision 2.0. 700 United States Environmental Protection Agency. (1993c). Determination of Total Kieldahl 701 Nitrogen by Semi-Automated Colorimetry Method 351.2, Revision 2.0. 702 Van de Bogert, M. C., Bade, D. L., Carpenter, S. R., Cole, J. J., Pace, M. L., Hanson, P. C., & 703 Langman, O. C. (2012). Spatial heterogeneity strongly affects estimates of ecosystem 704 metabolism in two north temperate lakes. Limnology and Oceanography, 57(6), 1689-705 1700. doi:10.4319/lo.2012.57.6.1689 706 Vilas, M. P., Marti, C. L., Adams, M. P., Oldham, C. E., & Hipsey, M. R. (2017). Invasive 707 Macrophytes Control the Spatial and Temporal Patterns of Temperature and Dissolved 708 Oxygen in a Shallow Lake: A Proposed Feedback Mechanism of Macrophyte Loss. 709 Frontiers in Plant Science, 8. doi:10.3389/fpls.2017.02097 710 Verpoorter, C., Kutser, T., Seekell, D. A., & Tranvik, L. J. (2014). A global inventory of lakes 711 based on high-resolution satellite imagery. Geophysical Research Letters, 41(18), 6396-6402. doi:10.1002/2014gl060641 712

- 713 Wilkinson, G. M., Carpenter, S. R., Cole, J. J., Pace, M. L., Batt, R. D., Buelo, C. D., & 714 Kurtzweil, J. T. (2018). Early warning signals precede cyanobacterial blooms in multiple 715 whole-lake experiments. Ecological Monographs, 88(2), 188-203. doi:10.1002/ecm.1286 716 Wu, X. D., Kong, F. X., Chen, Y. W., Qian, X., Zhang, L. J., Yu, Y., ... Xing, P. (2010). 717 Horizontal distribution and transport processes of bloom-forming Microcystis in a large 718 shallow lake (Taihu, China). *Limnologica*, 40(1), 8-15. doi:10.1016/j.limno.2009.02.001 719 Wynne, T. T., & Stumpf, R. P. (2015). Spatial and Temporal Patterns in the Seasonal 720 Distribution of Toxic Cyanobacteria in Western Lake Erie from 2002-2014. Toxins, 7(5), 721 1649-1663. doi:10.3390/toxins7051649 722 Zhou, Y. T., Obenour, D. R., Scavia, D., Johengen, T. H., & Michalak, A. M. (2013). Spatial and 723 Temporal Trends in Lake Erie Hypoxia, 1987-2007. Environmental Science &
- 724 *Technology*, 47(2), 899-905. doi:10.1021/es303401b

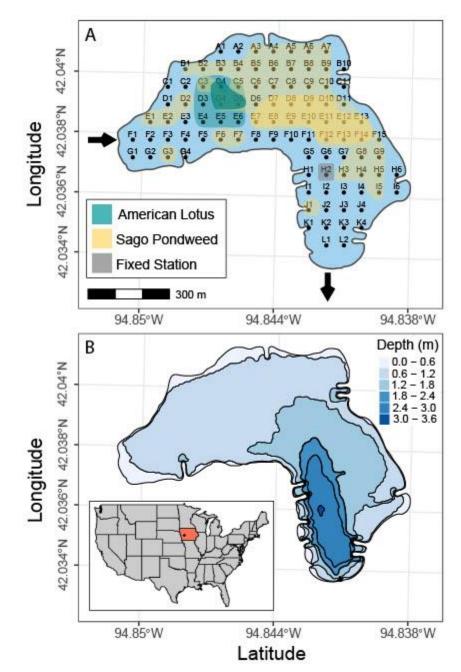
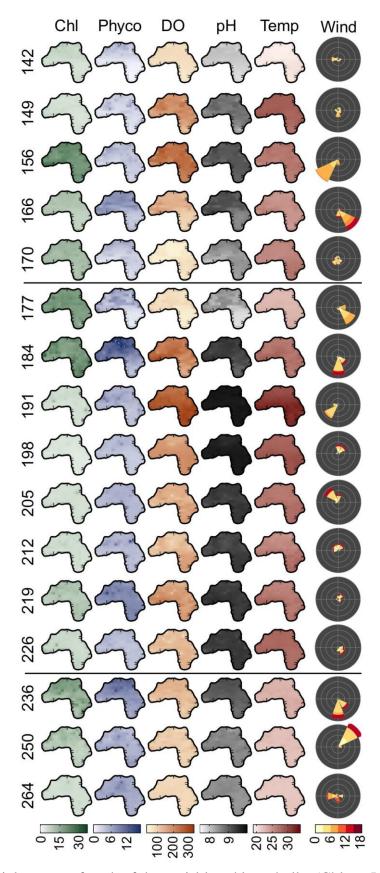


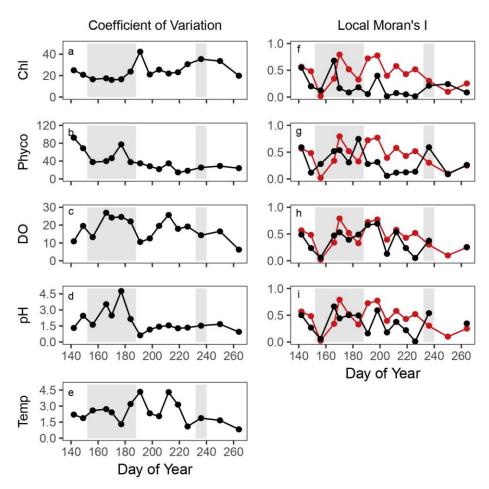
Figure 1. Sampling locations on a 65 m square grid of Swan Lake, a 40.5 hectare waterbody in western Iowa, USA. The main inlet to the lake and only outlet indicated with arrows. a) The location of the macrophyte beds of the two dominant species within the lake are shown on the map, with darker shading indicating the regions with the densest vegetation and the location of

- the high frequency sensor, b) the bathymetry of the lake and location of the lake in the state of
- 732 Iowa, in reference to the United States of America.



**Figure 2.** The spatial pattern of each of the variables chlorophyll *a* (Chl,  $\mu$ g L<sup>-1</sup>), phycocyanin

734 (Phyco, µg L<sup>-1</sup>), dissolved oxygen (DO, percent saturation), pH, and temperature (Temp, °C) for 735 each sampling event. The 98 sampling locations taken in a 65m grid (Figure 1) were interpolated 736 to a 25m grid using spatial inverse distance interpolation for visualization here. The color ramps 737 for each variable are scaled from the lowest to the highest value observed over the course of the 738 season across all sampling locations. The wind roses are the wind speeds (m s<sup>-1</sup>; color ramp) and 739 direction the wind came from for the 24 hours prior to a sampling event. The concentric circles 740 are the frequency of winds from that direction for the 24 hour period (expressed as a percentage, 741 largest circle is 80% of the time). In the case of a longer "spoke", the greater amount of time the 742 wind was from that direction. The horizontal lines between DOY 170 and 177, and DOY 226 743 and 236 mark the two large precipitation events that occurred between those sampling dates.

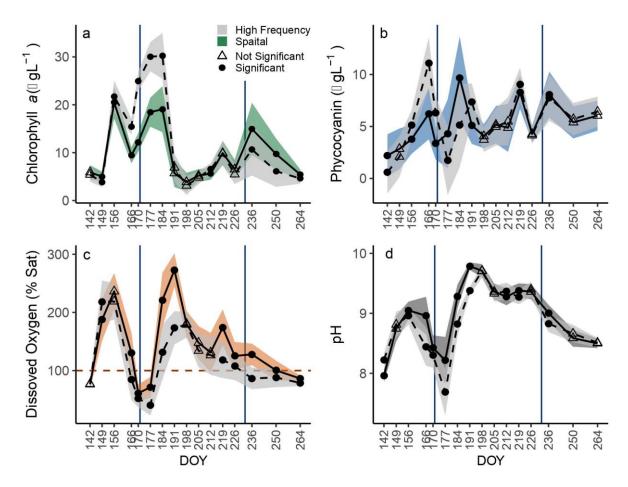


744 **Figure 3.** Time series of the spatial coefficient of variation (CV) and spatial autocorrelation (AC;

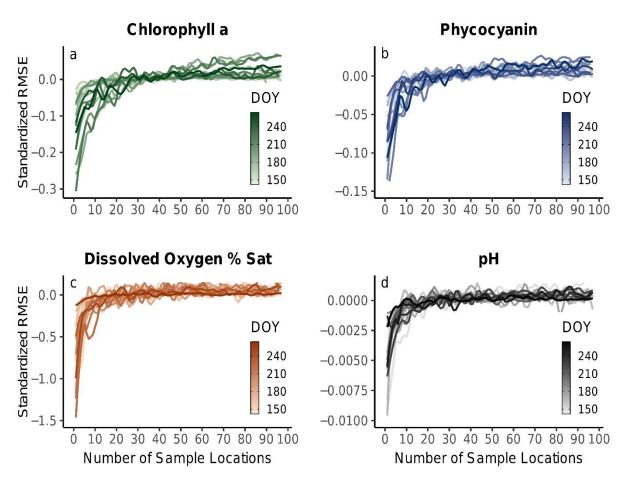
745 local Moran's I) of the biologically-mediated variables in Swan Lake (same variable

abbreviations as Figure 2). The gray polygons indicate periods of algal bloom. The red line is the

time series of temperature local Moran's I for comparison.



748 Figure 4. Comparison of the mean (lines and points) and range (shaded polygon) of 749 measurements from the spatial sampling and fixed station measurements. The fixed station data 750 were trimmed to the period that spatial sampling occurred. A filled circle is used for the 751 sampling dates when the means from the two sampling approaches were significantly different 752 (p<0.05), and an open triangle is used for the sampling dates when the mean of the two 753 approaches were not significantly different. The dark blue vertical lines indicate the dates of the 754 two major precipitation events and the red dashed line in panel c is at 100% dissolved oxygen 755 saturation.



756 **Figure 5.** Standardized root mean squared errors (RMSE) of rarefaction analysis. Fit lines

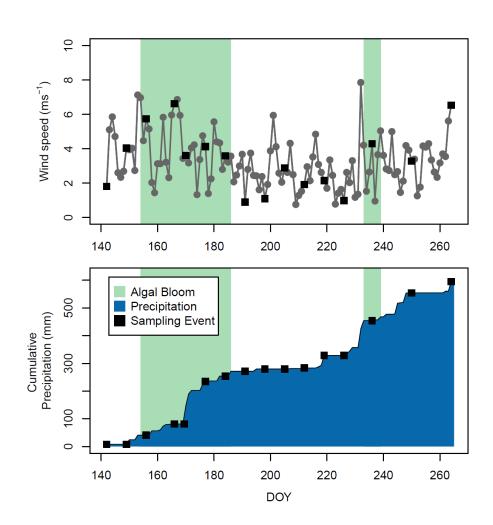
represent each sampling dates standardized RMSE (16 in total) and the gradient from light to

758 dark indicates first sampling event to last.

759	Supplementary material for
760	Title: Capturing the spatial variability of algal bloom development in a shallow temperate lake
761	Authors: Ortiz and Wilkinson
762	
763	Hourly Weather Data
764	Hourly wind and precipitation data were downloaded from the National Oceanic and
765	Atmospheric Automated Surface Observing System (NOAA ASOS) for Arthur N. Neu Airport
766	Carroll, Iowa, USA, less than 5km from the lake, through the Iowa State University Iowa
767	Environmental Mesonet (https://mesonet.agron.iastate.edu/). The data were summarized to daily

768 means and plotted as mean wind speed and cumulative daily precipitation.

## 769 Figures

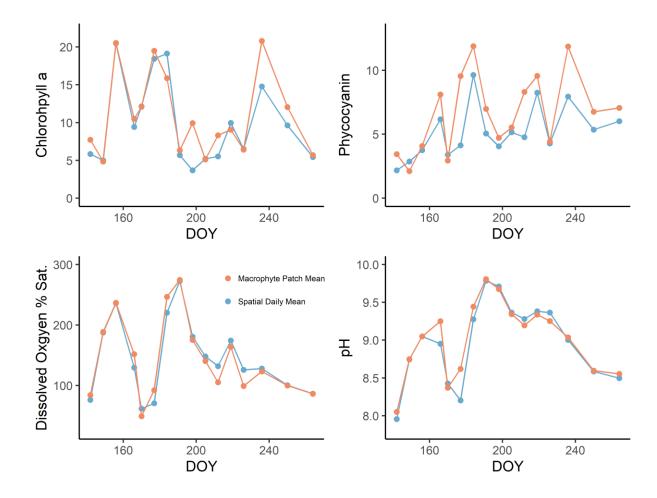


770

Supplement Figure 1. Daily meteorological data for the duration of the study. A) Daily mean
 wind speed during the study period. Periods of algal bloom in the lake are denoted by the green

polygons. B) The cumulative precipitation from the first to the last day of the study period. There

774 were two large precipitation events from DOY 170-171 and on DOY 132.



Supplement Figure 2. A comparison of the mean value for each variable in the American lotus
 macrophyte patch (sites C4, D4, D5) and the rest of the lake.