

1 **Stress Recovery for the Particle-in-cell Finite Element Method**

2

3 Authors: Haibin Yang¹, Louis N. Moresi^{1,2}, John Mansour³

4 Affiliation 1: School of Earth Sciences, University of Melbourne, Melbourne, Australia

5 Affiliation 2: Research School of Earth Sciences, Australian National University, Canberra, Australia

6 Affiliation 3: Monash eResearch Centre, Monash University, Clayton, Australia

7 Corresponding Author: Haibin Yang

8 Email: haibiny@student.unimelb.edu.au

9 *Non-peer reviewed preprint submitted to EarthArXiv*

10

11 **Abstract**

12 The interelement stress in the Finite Element Method is not continuous in nature, and stress
13 projections from quadrature points to mesh nodes often causes oscillations. The widely used
14 particle-in-cell method cannot avoid this issue and produces worse results when there are
15 mixing materials of large strength (e.g., viscosity in Stokes problems) contrast in one element.
16 The post-processing methods including (1) distance weighted average from surrounding
17 particles to the centroid mesh node (Post-local), (2) global projection with least square fit
18 (Post-global), and (3) superconvergent point recovery method (SPR), cannot effectively
19 eliminate the stress fluctuations. We propose three pre-processing methods to reduce the
20 interface contrast in mixing elements: (1) global method with harmonic-mean averaging
21 (GHM), (2) unification of properties at mixed-material elements (UnE), and (3) averaging
22 particle properties within a specified distance to gauss quadrature points (AGP). For tests of
23 Q_1 elements, the results processed by combining either pre-processing method with the
24 Post-local projection can increase the precision. The GHM pre-processing method is the least
25 computationally expensive application and the easiest to implement, the AGP pre-processing
26 method is the most expensive and the UnE in-between. However, for Q_2 elements, the GHM
27 pre-processing method fails in stress recovery, and produces worse results than those without
28 any pre-processing procedures. For general cases (both Q_1 and Q_2 elements), the AGP pre-
29 processing method is recommended. The optimal sampling radius used in the AGP method
30 is close to that size of one element, beyond which it increases computational time, but does
31 not significantly increases the accuracy of recovered stresses. In terms of the averaging

32 approaches used in the AGP method, the harmonic mean is suitable for simple-shear-
33 dominated processes and the arithmetic mean is better for the pure-shear-dominated
34 models. For complex models, the AGP method of harmonic mean combined with the SPR
35 post-procedure is recommended. The AGP method is found to be able to efficiently reduce
36 stress perturbations in a synthetic model of complex fault geometries like the San Andreas
37 Fault system.

38

39 **Key words:**

40 Particle-in-cell; Finite Element Method; Stress fluctuation; Stress smoothing; Numerical
41 geodynamic modelling

42

1 Introduction

43

44 The classical Finite Element Method (FEM) (see e.g. [Hughes \(2012\)](#)) has been widely used to
45 simulate different structures in engineering. Different from most engineering problems,
46 geological simulations are challenged by emergent structures due to the non-linear
47 processes involved ([Lenardic et al., 2003](#)). For the body fitted FEM meshes, the modelled
48 system evolving with large deformations distorts meshes, which may produce more
49 complexities (like re-meshing) in computation ([Braun and Sambridge, 1994](#)). Alternatively, the
50 particle-in-cell (PIC) method allows Lagrangian material particles to move in a background
51 Eulerian mesh ([Harlow, 1964](#); [Sulsky et al., 1994](#)). Those particles carry information of density,
52 composition, viscosity, etc., while the unknowns are solved at nodes of the mesh.

53

54 One disadvantage of the PIC FEM is that large deformation causing mixing of materials
55 of great strength contrast in one cell triggers oscillations of stresses along the cell edges. For
56 Stokes problems, this strength property generally refers to viscosity. The jump of viscosity in
57 one element is found to give rise to an error more than two orders of magnitude larger than
58 ones with material interfaces aligned with element boundaries ([Moresi et al., 1996](#)). This
59 viscosity jump can also significantly degrade convergence rate of numerical solvers ([May and](#)
60 [Moresi, 2008](#)).

61

62 To avoid this mixing effect, [Wallstedt and Guilkey \(2011\)](#) suggested a weighted least
63 square particle-in-cell method to project materials properties to nodes, which requires sub-

64 division of the mixed-material cell. They integrate those particles of same properties
65 separately. To save the computational time of dividing the integration domain, [Sadeghirad et](#)
66 [al. \(2011\)](#) tried new basis functions over particle field using a 4-node finite element
67 interpolation method and the integration is based on each particle in the corresponding
68 element rather than on Gauss quadrature points, which can be implemented in the mantle
69 convection models in a simple and efficient way ([Moresi et al., 2003](#)). While this may work
70 well for simple 2D problems, its efficiency of implementation for complex geological problems
71 is not sufficiently studied.

72

73 There are some other, simpler methods, which directly manipulate the property
74 distribution in one element. Averaging of the viscosity in cells with multi-phase materials in
75 the Finite Difference Method is systematically studied by [Deubelbeiss and Kaus \(2008\)](#).
76 However, for the case of the Finite Element Method, they simply put a single value for all
77 quadrature points in one element to avoid viscosity jumps. This constant interpolation of
78 viscosity in the cell with mixed materials is also tested by [Thielmann et al. \(2014\)](#), who
79 proposed another more sophisticated strategy, the linear least square interpolation, which
80 requires solving a linear equation for each mixed-material element. However, this extra cost
81 of combining linear least square interpolation with Q_2 element does not work well in cases of
82 sharp changes of viscosity in one element ([Thielmann et al., 2014](#)).

83

84 This work examines several smoothing methods to eliminate spurious stress fluctuations

85 in the framework of Underworld2, which is a Python application programming interface to
86 simulate geodynamics processes (<https://github.com/underworldcode/underworld2>) ([Moresi](#)
87 [et al., 2007](#)). The smoothing methods we test can be conveniently implemented and efficiently
88 run with Underworld2. We first introduce two post-processing methods available in
89 Underworld2, and then describe three pre-processing solutions: (1) a global method with
90 harmonic-mean averaging, (2) unifying material properties at elements with mixed materials,
91 and (3) averaging at gauss quadrature points, which, in the remainder of this context, are
92 referred to as GHM, UnE, and AGP respectively. Additionally, we further compare them with
93 the classical stress recovery technique that reconstruct continuous stresses on specified
94 patches based on super-convergent points inside the elements (SPR) ([Zienkiewicz and Zhu,](#)
95 [1992a, b](#)). The SPR method is for cases without the internal structure in the element, so they
96 are not intended to alleviate the problem caused by mixed-material elements. The effect of
97 each method is checked with geological models that have analytical solutions. Combining
98 effects of different pre- and post-processing methods are also present. Finally, we investigate
99 a synthetic model with relatively complex fault geometries based on the fault data from the
100 San Andreas Fault system.

101

102

2 Governing equations

103 The simulation is based on the Stokes equation for Newtonian viscous, incompressible flow:

$$\frac{\partial \sigma_{ij}}{\partial x_j} - \frac{\partial P}{\partial x_i} = \rho g_i \quad (1)$$

$$\frac{\partial v_i}{\partial x_i} = 0 \quad (2)$$

$$\dot{\sigma}_{ij} = 2\eta \dot{\varepsilon}_{ij} \quad (3)$$

$$\dot{\varepsilon}_{ij} = \frac{1}{2} \left(\frac{\partial v_i}{\partial x_j} + \frac{\partial v_j}{\partial x_i} \right) \quad (4)$$

104 where σ_{ij} denotes deviatoric stress, P pressure, ρ density, g_i gravity acceleration, v_i
 105 velocity, η Newtonian viscosity, ε_{ij} strain rate, and the Einstein summation convention is
 106 used here.

107

3 Methods

108

3.1 Post-processing methods

109 For the classical Finite Element Method, inter-element stress continuity is not guaranteed.
 110 Smoothing the numerical discontinuity to fit the physically continuous system has been
 111 conventionally implemented as a post processing step. These generally include (1) averaging
 112 around local nodes and (2) global projection with least square fits. In this study, they are taken
 113 as the Post-local and Post-global projections, respectively ([Hinton and Campbell, 1974](#)).

114

3.1.1 Post-local - averaging around local nodes

115 The projection at mesh nodes is done through extrapolation from superconvergent points

116 (e.g., Gauss points) and then averaging locally at mesh nodes. In the FEM framework, the
 117 distance-based weight can be achieved through the shape function (N_a) in the whole
 118 calculation domain Ω , so the average nodal stress

119

$$\sigma^h = \frac{\int N_a \sigma^p d\Omega}{\int N_a d\Omega} \quad (5)$$

120 where σ^p is the point stress. Note that, although the integration is applied to the whole
 121 domain, the shape function is zero outside adjacent elements.

122

123 3.1.2 Post-global - global projection with least square fits

124 The error between the smoothed stresses $\tilde{\sigma}$ and the point stresses σ^p is

$$\int (\tilde{\sigma} - \sigma_p)^2 d\Omega \quad (6)$$

125 The smoothed stress is written as

$$\tilde{\sigma} = \sum_{B=1}^{n_{np}} N_B \sigma_B^h \quad (7)$$

126 where n_{np} is the number of nodal points. The least square method is used to minimize the
 127 error by

$$\frac{\partial}{\partial \sigma_B^h} \int (\tilde{\sigma} - \sigma_p)^2 d\Omega = 0 \quad (8)$$

128 for $B = 1, 2, \dots, n_{np}$. This yields the following matrix problem:

$$\mathbf{X}\tilde{\boldsymbol{\sigma}} = \mathbf{P} \quad (9)$$

129 where $\mathbf{X} = [X_{AB}]$, $\tilde{\boldsymbol{\sigma}} = \langle \sigma_B^h \rangle$ and $\mathbf{P} = \langle P_A \rangle$.

130 The construction of \mathbf{X} and \mathbf{P} is implemented in the elementwise fashion:

$$x_{ab}^e = \int_{\Omega_e} N_a^e N_b^e d\Omega, \quad p_a^e = \int_{\Omega_e} \sigma_p N_a^e d\Omega \quad (10)$$

131 for $1 \leq a, b \leq n_{en}$, where n_{en} is the number of nodal points per element.

132

133 The global projection method is a more costly process to recover accurate nodal stresses and

134 sometimes produces overshoot values at nodal points.

135

136 3.1.3 SPR-superconvergent point recovery

137 The SPR method ([Zienkiewicz and Zhu, 1992a](#)) compute a continuous nodal stress field, σ^* ,

138 from a patch of elements local to each node:

$$\sigma^* = \mathbf{M} \mathbf{a} \quad (11)$$

139 where $\mathbf{M} = [\mathbf{1}, x, x^2, \dots, x^m]$ and $\mathbf{a} = [a_1, a_2, a_3, \dots, a_{m+1}]^T$ for one dimensional cases.

140 m is the order of the polynomial expansion, which is the same as that used in the shape

141 function N . Accordingly, for the two-dimensional expansion for linear elements, $\mathbf{M} =$

142 $[1, x, y]$, and for quadratic $\mathbf{M} = [1, x, y, x^2, xy, y^2]$ (the xy term is optional).

143 To determine the unknown parameters \mathbf{a} in equation (11), we minimize

$$\mathbf{F} = \sum_{i=1}^n (\sigma_i^h(x_i, y_i) - \mathbf{M}(x_i, y_i) \mathbf{a})^2 \quad (12)$$

144 where (x_i, y_i) are the coordinates of specified sampling points, the total number of which is

145 n . Therefore, for \mathbf{F} to be a minimum

$$\frac{\partial \mathbf{F}}{\partial \mathbf{a}_i} = 0 \quad (13)$$

146 This gives

$$\sum_{i=1}^n \mathbf{M}^T(x_i, y_i) \mathbf{M}(x_i, y_i) \mathbf{a} = \sum_{i=1}^n \mathbf{M}^T(x_i, y_i) \sigma^h(x_i, y_i) \quad (14)$$

147 This system is rewritten as

$$\mathbf{a} = \mathbf{A}^{-1} \mathbf{b} \quad (15)$$

148

149 where

$$\mathbf{A} = \sum_{i=1}^n \mathbf{M}^T(x_i, y_i) \mathbf{M}(x_i, y_i) \text{ and } \mathbf{b} = \sum_{i=1}^n \mathbf{M}^T(x_i, y_i) \sigma^h(x_i, y_i) \quad (16)$$

150 After obtaining the parameter \mathbf{a} in equation (11), with the polynomial expansion, the nodal

151 values can be evaluated by any given coordinates to the functional form of σ^* in equation

152 (11).

153 3.2 Pre-processing method

154 The classical post-processing methods are designed to produce continuous, node-based

155 stress field, but not to resolve the stress perturbations caused by mixed-material elements.

156 Instead, we utilize pre-processing methods to reduce the strength contrast across the

157 interfaces between two materials. The Post-local method is taken as the default post-

158 processing procedure that projects stress from particles to nodes after one specific pre-

159 processing method. Any other post-processing method is stated when used in this study.

160 3.2.1 GHM - global method with harmonic-mean averaging

161 We use the method described in section 3.1.1 to project reciprocals of the property

162 information η^p (e.g., viscosity) carried by particles to mesh nodes through the shape function
 163 (N_a) in the whole calculation domain Ω , so the harmonic mean viscosity

$$\eta^h = \frac{1}{\frac{\int N_a (1/\eta^p) d\Omega}{\int N_a d\Omega}} \quad (17)$$

164 The strength values η^* considered in the elementwise integration is

$$\eta^* = \sum_{A=1}^{n_{np}} N_A \eta_A^h \quad (18)$$

165 **3.2.2 UnE - unification at one element**

166 Here we first look for elements that contain more than one material, and then unify the
 167 properties of all the particles in those elements to be one value. The harmonic mean method
 168 averaging over all types of points in one cell, gives the unified property value

169

$$\eta^e = n_k / \sum_{A=1}^{n_k} 1/\eta_A^p \quad (19)$$

170 where n_k is the number of material types in one element. It is worthwhile noting that n_k is
 171 not the number of particles in one element. For large viscosity jump $> 10^3$ Pa · s, if the particles
 172 numbers in one cell is few (e.g. 10s), the effective viscosity for both cases are the same order
 173 of magnitude. For simplicity, we take the number of material types. The averaging over the
 174 number of particles is introduced in the next section.

175 **3.2.3 AGP-averaging at Gauss quadrature points**

176 This method is also applied only to elements with a mixture of materials, and we first locate

177 mixed-material cells. Then, instead of unifying materials over one cell, particles within distance
178 δ to the selected Gauss quadrature point are averaged by

$$\eta^{gp} = n_{dp} / \sum_{A=1}^{n_{dp}} 1/\eta_A^p \quad (20)$$

179 where n_{dp} is the number of sampling points for the specified Gauss quadrature point. The
180 sampling numbers n_{dp} is determined by the selected distance δ . The effect of sampling
181 distance is discussed later. The default averaging method is harmonic mean as shown in
182 equation (21), and other averaging methods are stated when used.

183

184 **4 Results and discussion**

185 **4.1 Models with analytical solutions**

186 **4.1.1 Simple shear model**

187 We first test the effect of different smoothing methods by a simple shear model. Details of
188 the setup is shown in [Fig. 1a](#). A narrow weak zone ($\eta_2 = 10^{18} \text{ Pa}\cdot\text{s}$) is embedded in the
189 background of high viscosity ($\eta_1 = 10^{20} \text{ Pa}\cdot\text{s}$). With the top boundary fixed, the driving
190 velocity V_0 at bottom is 1 cm yr^{-1} , and periodic boundary conditions are applied to the left
191 and right boundary. The model size is $20 \text{ km} \times 20 \text{ km}$. Rectangular meshes with Q_1 elements
192 of 40×40 , 80×80 , 160×160 , and 320×320 are compared. The thickness of the weak zone
193 is $h_2 = 2 \times 1.9 \text{ dy}$, where dy is one element length in y direction for the element resolution

194 of 40×40 , and the thickness of the weak zone does not change with mesh resolution. The
195 weak zone is centered at $y = 0$, so neither interface (at $\pm 1.9 \, dy$) is aligned with element edges.
196 Taking the solution of Couette flow and considering stress continuity at both interfaces, we
197 give the analytical results for shear stress

$$\sigma_{xy} = \frac{V_0 \eta_1}{h_3 + h_1 \left(\frac{h_2 \eta_1}{h_1 \eta_2} + 1 \right)} \quad (21)$$

198 The results of 80×80 element resolution are shown in Fig. 1b-1f. The simple shear model
199 should produce a homogeneous stress field, and the exact value from the analytical model is
200 $10^{5.18}$ Pa. We define the normalized relative error

$$\text{err} = \frac{f - f_0}{f_0} \quad (22)$$

201

202 where f is the nodal value from numerical models and f_0 is the corresponding analytical
203 solution at the same point. In this simple shear model, $f_0 = 10^{5.18}$ Pa. The mixing effect is
204 obvious in the default output with Post-local projection (averaging around local nodes) (Fig.
205 1b), which results in fluctuations in stress error from -24.6 % to 78.0 %. We note that the
206 fluctuations in Post-local projections are interpolated results from the raw values in particles,
207 as the equation (5) does. We check the stresses at particles and found the same spurious
208 stresses along interfaces (not shown in figures). Directly applying the superconvergent point
209 recovery (SPR) method on the Post-local stress field, the resulted relative error range is
210 reduced to 9.9 % – 36.7 % (Fig. 1c). We implement the pre-processing method to degrade
211 stress perturbations. For the GHM pre-processing method, the relative error is -10.4 % – 28.1 %

212 (Fig. 1e). Further applying the SPR post-processing method on the GHM results, we find the
 213 relative error range is narrowed to -4.0 % – 14.5 % (Fig. 1f). In contrast, the UnE pre-processing
 214 method leads to a nearly homogeneous stress field as we expect, but it is about 8 % higher
 215 than the analytical solution (Fig. 1d).

216

217 To have a better understanding of the resolution controls on the numerical results
 218 accuracy, we define the mean error

$$\text{Err} = \frac{\sqrt{\sum_{i=1}^n \text{err}^2}}{n} \quad (23)$$

219

220 where n is the number of nodal values in a specified area. We select nodes in a rectangular
 221 domain $-3.4 dy < y < 3.4 dy$, which covers the area of intensive oscillations. The
 222 resolution test (Fig. 2a) demonstrates that the errors decrease by three orders of magnitude
 223 for all applied smoothing methods. Additionally, we find the Post-global projection (global
 224 projection with least square fits) has slightly lower accuracy than the Post-local (averaging
 225 around local nodes) method. The errors in both the GHM and UnE pre-processing methods
 226 are about one order of magnitude less than Post-local projection results. Additionally, we
 227 define a parameter to describe the relative fluctuation range

$$flt = \frac{\max(\mathbf{f}) - \min(\mathbf{f})}{\text{mean}(\mathbf{f})} \quad (24)$$

228 where \mathbf{f} is a group of all nodal values in a specified domain. The fluctuation ranges do not
 229 decrease dramatically (Fig. 2b) as the error does, while the grid resolution increases from 40
 230 to 320. Using the SPR post-processing method to further process GHM results can help
 231 reduce fluctuation ranges by 10 %. Generally, the UnE post-processing method produces a
 232 much narrower fluctuation range than other results, because we unify the viscosity values in
 233 those mixed-material elements in the simple shear model. We note that the high efficiency

234 of the UnE method is limited to special cases, like the simple shear model with material
235 interfaces parallel to shear directions.

236

237 For the AGP pre-processing method (averaging at Gauss quadrature points), we test
238 several models with different δ (the sampling radius of the circle centering at Gauss points
239 in mixed-material elements). Fig. 3 shows results of a model (80×80) after the AGP pre-
240 processing procedure. With sampling radius increasing from $0.5dx$ to $4dx$, where dx is one
241 element length, the relative error ranges of stress field perturbations decrease from $-8.4\% -$
242 22.9% to $4.4\% - 9.9\%$. The perturbation ranges should reduce with further increasing sampling
243 radius, since, in the extreme case of a global scale averaging, it nearly produces a unified
244 property in those mixing elements, thus yielding the same effect as the UnE pre-processing
245 method. We note that the computation time for the UnE pre-processing method also
246 increases with search area.

247

248 Q_2 elements are also tested to investigate how efficiently each method is in reducing
249 stress perturbations (Fig. 4), which are more severe for Q_2 elements than the Q_1 cases. The
250 Post-local projection produces relative errors of $-99.6\% - 243.6\%$ (Fig. 4a). The GHM method
251 yields the worst result (err = $-99.8\% - 3047.5\%$; Fig. 4b), which is markedly worse than the
252 Post-local projection. The UnE method can still generate uniform stress field (Fig. 4c), and the
253 relative error ($\sim 8\%$) is close to that in Q_1 cases in Fig 1d. The stress processed by the AGP
254 method with sampling radius of $1dx$ dampens most of the noise to a much lower level (-0.2%
255 $- 7.5\%$) than the Post-local and GHM methods (Fig. 4d). In the next step, the SPR method is
256 utilized to do further post-processing of those results from Fig. 4a-4d. The SPR method
257 reduces stress perturbations by more one order of magnitude for the Post-local solution (Fig.
258 4e). Although upper bound of the error after applying SPR on GHM results degrades by about
259 one order of magnitude, the error ranges still have more than two orders of magnitude in
260 spurious stress oscillation (Fig. 4f).

4.1.2 Folding model

261
262 Although the simple shear model has illustrated some features of different methods, models
263 with a relatively complex geometry and boundary conditions may help us learn more about
264 the broader applicability of those methods. We use the folding model from [Ramberg \(1962\)](#)
265 as a benchmark. The model investigates contact strain in an incompetent host rocks
266 adjacent to a buckling thin sheet that is composed of competent rocks. Mathematically, the
267 competent buckling rock is treated as a very viscous Newtonian material (10^{22} Pa·s), while
268 the background incompetent rock is of lower viscosity (10^{18} Pa·s) than the embedded
269 rock. The geometry of the buckling sheet is described by a sinusoidal function

$$y = y_0 \sin \frac{2\pi x}{\lambda} \quad (25)$$

270 where y is the deflection of the center line of the thin sheet, y_0 the amplitude and λ the
271 wavelength ([Fig. 5](#)). In the physical analysis, the thickness of the thin layer is assumed to be
272 extremely small relative to its wavelength. We simulate a domain of 20 km × 20 km with the
273 Q₁ element resolution of 320 × 320. The thickness of the thin layer (0.125 km) spans 2 grid
274 size (3 grid points), and the wavelength is 2.5 km, which is 1/8 of the model width. The
275 bottom boundary is designed to follow the sinusoidal shape, and y_0 in equation (25) takes
276 half grid size. In this case, all other elements have rectangular shape except the lowest row
277 at the base. The interface between two materials crosses elements ([Fig. 5](#)). The high
278 resolution is required to refine the sinusoidal geometry to be comparable with analytical
279 solutions. As the relative fluctuation range does not vary too much with model resolution
280 ([Fig. 2b](#)), the high-resolution models still have the issue of stress perturbations.

281

282 Free slip boundary conditions are applied to all boundaries except the bottom, where
283 vertical velocity is prescribed

$$V_y = V_0 \sin \frac{2\pi x}{\lambda} \quad (26)$$

284 where $V_0 = 0.5 \text{ cm yr}^{-1}$. With the stream function method, [Ramberg \(1962\)](#) gave the
285 functional form for the shear stress

$$\sigma_{xy} = 2V_0 \omega^2 y e^{-\omega y} \cos(\omega x) \quad (27)$$

286 where $\omega = \frac{2\pi}{\lambda}$.

287

288 Comparing the numerical results with analytical results at $y = 3dy$, where dy is the spatial
289 resolution in y direction, [Fig. 6](#) shows the results processed by the Post-local projection, GHM,
290 UnE and AGP methods. The worst results happen at the situation of pure Post-local without
291 any pre-processing procedure, and several orders of magnitude differences between the
292 analytical and numerical solutions occur around where the peak value is ([Fig. 6a](#)). Generally,
293 all other cases with the application of a pre-processing method have damped the stress misfit
294 error at $n\lambda \leq x \leq (n + 1/2)\lambda$, where n is integer ([Fig. 6b, c & d](#)). For the pre-processing
295 methods, the maximum error for the UnE method is 31%, and the GHM method is 17%. The
296 best one comes from the AGP method where the misfit at peak values are less than 0.4%.

297

298 Particularly, we further test the best choice δ used in AGP. The sampling radius used in
299 [Fig. 6d](#) is equal to one element size. Comparing with other choices of sampling radius ([Fig. 7](#)),

300 we find the maximum misfit between numerical results and analytical solutions quickly
301 decreases from half to one element length for both the observation at $y = 3dy$ and $4.5dy$, and
302 the misfit slightly varies with the sampling radius when it is larger than one element size. For
303 the observation at $y=3dy$, It is consistent with our conclusion in section 4.1 that the end-
304 member case of the AGP method that has an extremely large sampling area produces the
305 same effect of the UnE method (Fig. 6b), which generates worse results than the AGP method
306 with $\delta/dx \cong 1$ (Fig. 6d).

307

308 This trend of maximum misfit versus sampling radius in models with Q_2 elements is
309 consistent to that observed in models with Q_1 elements (Fig. 7). The mesh has 160×160 Q_2
310 elements. The thickness of the thin layer (0.25km) also spans 2 grid size (5 grid points), and
311 the wavelength is 2.5 km, which is $1/8$ of the model width. The maximum misfit drops from
312 more than 1×10^4 Pa to less than 7×10^3 Pa when the sampling radius increases from half to
313 one element length for both observations at $y = 3dy$ and $4.5dy$, after which the maximum
314 misfit does not vary significantly with sampling areas. In this case, the optimal sampling radius
315 for AGP method is around one element length, which is same as that for Q_1 elements.

316

317 Additionally, two other averaging methods, the arithmetic mean and geometric mean,
318 are tested in the AGP method instead of the harmonic mean used in equation (21). For a
319 positive sequence $\eta = (\eta_1, \eta_2, \dots, \eta_n)$, the arithmetic mean (An) and the Geometric mean
320 (Gn) are defined respectively by

$$An(\eta) = \frac{1}{n} \sum_{i=1}^n \eta_i \text{ and } Gn(\eta) = \frac{1}{n} \prod_{i=1}^n \eta_i \quad (28)$$

321 The harmonic mean defined in equation (20) is referred to as Hn . It is well known and has
322 been demonstrated by [Xia et al. \(1999\)](#) that for the same sequence, $An \geq Gn \geq Hn$. The
323 misfit of AGP method with An and Gn yields a maximum misfit (not shown in figures) that
324 is close to that with the default Post-local projection ([Fig. 6a](#)), which has 2-3 orders of
325 magnitude in difference between the numerical and analytical solutions. This indicates that
326 the harmonic mean averaging method (Hn), which gives more weight to smaller values in the
327 selected dataset, may be preferred to describe the interface between two materials of great
328 contrast in strength.

329

330 **4.1.3 SolCx model**

331 The SolCx benchmark model is used by [Thielmann et al. \(2014\)](#) to test the influence of sharp
332 viscosity jump within one element on computation error and convergence rate. The SolCx
333 benchmark is a complementary case that includes the body force terms while that is not
334 considered for both simple shear and folding models. We use the model setup from
335 [Thielmann et al. \(2014\)](#) to compare the AGP pre-processing method with their sophisticated
336 least square interpolation method, which is referred to as T2014 model in the next context.
337 The analytic solution is derived by ([Zhong, 1996](#)), and the code for this analytic solution is
338 included in the Underworld package.

339

340 The model is a unit box with a viscosity jump of 10^3 at $x=0.5$, and density is described by
341 the trigonometric function $\rho = \sin(\pi y)\cos(\pi x)$, and the gravity is 0 in the x direction and 1
342 in the y direction (Fig. 8). The mesh is composed of 51×51 Q_2 elements, forming 103×103
343 calculation nodes. The AGP pre-processing method with the sampling radius of one element
344 length is applied to compare with T2014 model results. The Harmonic-mean method is
345 utilized to average particle properties to gauss point. Free-slip boundary conditions are
346 applied to four edges.

347

348 The numerical results of velocity, shear stress and pressure field are shown in Fig. 9a-9d.
349 The corresponding error in Fig. 9e-9h is the absolute difference between the analytic and
350 numeric solutions. The maximum absolute error in our results is 2×10^{-5} for V_x and 2.5×10^{-4}
351 for V_y , which is much lower than that in the T2014 model (1.4×10^{-4} for V_x and 5×10^{-4} for V_y).
352 However, the error for pressure field in our model is about twice of the T2014 model. The
353 maximum shear stress error (8.8×10^{-3}) occurs to the interface in our model, but T2014 model
354 does not provide their stress data for comparison. The T2014 model uses least square
355 interpolation which requires the calculation of the inverse of a matrix, same as the SPR or the
356 Post-global method. Enforcing continuity through the least squares fit may introduce the
357 over-shooting or under-shooting issues, which produce properties at quadrature points or
358 nodes of significantly large (or small) values relative to the maximum (or minimum) value of
359 the physical properties on particles (Thielmann et al., 2014). Therefore, the AGP pre-
360 processing method is a much more straightforward and robust implementation than the least

361 square interpolation used in the T2014 model.

362

363 Instead of the harmonic-mean averaging method, with the arithmetic-mean averaging,
364 we find the maximum shear stress error is reduced to 5×10^{-3} (Fig. 10). In this case, the
365 arithmetic-mean may be more suitable than the harmonic-mean averaging method. Shear
366 stress errors at $x=0.51$ for different methods are illustrated in Fig. 10c. The difference caused
367 by arithmetic and harmonic mean may be explained by their physical meanings proposed by
368 [Schmeling et al. \(2008\)](#), who suggested that the harmonic-mean averaging best represents
369 the effective viscosity of simple shear models, while the arithmetic-mean averaging can stand
370 for the effective viscosity of pure shear models. The SolCx model is driven by body force
371 rather than the surface drivers in simple shear models and folding models. Accordingly, the
372 SolCx model may be dominated by the pure-shear mode and the other two models in section
373 4.1 and 4.2 are by the simple shear mode. In real cases, complex geological models are often
374 composed of both end-member cases. Thus, the choice of arithmetic or harmonic mean
375 depends on the specific geologic problems. However, we further apply the SPR post-
376 processing method on those results obtained by the AGP pre-processing method with
377 harmonic mean averaging, and the final accuracy can reach the same level as the AGP with
378 arithmetic-mean method (Fig. 10). Therefore, the AGP pre-processing with harmonic
379 averaging plus the SPR post-processing method could be a practical option for models when
380 we are not familiar with the dominating deformation mode (simple shear or pure shear).
381 Alternatively, it is worthwhile to consider the transversely isotropic viscosity that takes different

382 averaging method for the viscosity in different directions in the future work.

383 4.2 Models without analytical solutions

384 Models in sections 4.1 are all based on a relatively simple geometry, the stress field of which
385 can be obtained through analytical solutions, and the precision of numerical results have been
386 checked by plotting against corresponding analytical solutions (Fig. 6 & 9). In this section, we
387 test models that do not have a relatively simple analytical solutions as those in section 4.1.
388 First, we test a model with a fault (thin weak zone) that is at a low angle to the x-axis. Then,
389 we build a synthetic model by using the observations from a complex fault network to test
390 how the methods proposed in this study work in a complex system that does not have
391 analytical solutions.

392 4.2.1 Fault at a low angle to x-axis

393 The model tested here is close to that investigated in section 4.1.1 and has a weak zone ($\eta_2 =$
394 $10^{18} Pa \cdot s$) embedded in a relatively strong background ($\eta_1 = 10^{20} Pa \cdot s$). The setup is
395 shown in Fig. 11. With the top boundary fixed, the driving velocity V_0 at bottom is 1 cm yr^{-1} .
396 The velocity at both lateral boundaries linearly increases from 0 cm yr^{-1} (top boundary) to 1
397 cm yr^{-1} (base boundary). The fault ($\sim 160 \text{ m}$ thick) is at $\sim 10^\circ$ to the x-axis, thus containing
398 rectangular elements that has both strong and weak materials (Fig. 11a). We also build a
399 model that has mesh edges aligned with the material interfaces (Fig. 11b) and the mesh in
400 the y-direction surrounding the weak zone is refined, with the grid size varying from 0.21 m
401 to 330 m . We take the aligned case as a reference model and the results may represent

402 analytical solutions due to high resolution and body fitted mesh.

403

404 The reference model demonstrates smooth stress field along the fault (Fig. 11c), while
405 the default Post-local projection has intensive perturbations along the fault (Fig. 12). Besides
406 the absolute value of the stress field (left panel in Fig. 12), we also calculate the relative
407 differences between the reference model and the case that contain elements with a mix of
408 materials (right panel in Fig. 12). All pre-processing methods can help mitigate stress
409 perturbations as that in section 4.1. The relative difference between the Post-local projection
410 without any pre-processing and the reference model is high (0.6-1.0) along the fault. After
411 adjusting material contrast in elements with pre-processing methods, the difference is
412 reduced to be 0-0.2, though some light spots (high difference) occur along the fault. In this
413 case, the global method with harmonic-mean averaging (GHM) seems to produce better
414 results than averaging at Gauss quadrature points (AGP). For those light spot, further
415 application of the superconvergent point recovery method (SPR) may help alleviate it as
416 shown in section 4.1. It is worthwhile noting that, with the SPR method, we can reconstruct
417 continuous stresses on specified patches rather than on the whole domain. It is convenient to
418 use the patch recovery method on specified areas of anormal stress. In cases of 150×75 Q_2
419 elements, the results with the pre-processing procedure are not better than the Q_1 cases (not
420 shown in figures). When the yield stress (e.g., the upper limit of the stress is a constant value
421 of $10^{6.1}$ Pa) is applied, it is found to produce a smooth stress field. That means those high
422 stress perturbations along the fault are forced to be damped to be less than the yield stress,

423 but this will cause another problem that regions may yield when the expected stress is lower
424 than the yield value, which may affect fault branches development in the long-term
425 deformation.

426 **4.2.2 Complex fault geometries**

427 The case study of a complex fault geometry is based on the San Andreas Fault system in
428 California. The San Andreas Fault (SAF) System has received the most in-depth study around
429 the world, and scientific publications about the San Andreas Fault System from 1991-2013
430 produced about 3400 peer-reviewed articles ([Gizzi, 2015](#)), i.e., almost a paper is published
431 every other day. The San Andreas Fault strikes through the state of California and is a
432 transform boundary between the Pacific and North American plates. Many studies have
433 suggested that the San Andreas Fault at different segments accommodates 20-75 % relative
434 motion between these two plates ($\sim 50 \text{ mm yr}^{-1}$) ([Atwater and Stock, 1998](#); [DeMets and Dixon,](#)
435 [1999](#); [Meade and Hager, 2005](#)). We map the major faults based on the WGCEP (2007 Working
436 group on California Earth-quake Probabilities) fault traces, where only major fault traces with
437 long-term strain rate of orders of magnitude higher than that of less deformed areas are
438 selected ([Bird, 2009](#)) ([Fig. 11](#)). The 2-D model is 1110 km long and 484 km wide with the Q_1
439 element a resolution of 400×200 . A simple shear boundary condition is applied to $y = 0 \text{ km}$
440 ($V_x = 4 \text{ cm yr}^{-1}$) and $y = 484 \text{ km}$ ($V_x = 0 \text{ cm yr}^{-1}$), and the velocities at two lateral boundaries
441 linearly decrease from 4 cm yr^{-1} at $y = 0 \text{ km}$ to 0 cm yr^{-1} at $y = 484 \text{ km}$. The fault area generally
442 spans 3 – 5 grid points, and the viscosity for the mapped fault area is set to be a constant

443 value of 10^{19} Pa•s. The viscosity of non-fault area is 10^{23} Pa•s.

444

445 The second invariant stress of the viscous model in [Fig. 13](#) shows significant differences
446 between the raw result and the one with the AGP pre-processing method with the sampling
447 radius of one element length. Stress perturbations occur along the San Andreas Fault (SAF, x
448 = 250 – 500 km) and the Garlock Fault (GF; y = 150 km – 400 km) for the raw output with
449 default Post-local projection. With the AGP pre-processing method applied, a smooth stress
450 field is observed along the SAF and GF. The stress perturbation along the San Andreas Fault
451 and Garlock Fault in the raw output is one to two orders of magnitude higher than that the
452 smooth results. Although the analytical solution is unable to be obtained here for the complex
453 fault system, with benchmarks in section 4.1, we suggest that the results processed with AGP
454 pre-processing method is supposed to have higher precision than the raw result. From the
455 AGP result, we find that the areas surrounding fault tips are of high stress, while the areas
456 close to and along the major fault traces (the San Andreas Fault or Garlock Fault) are of low
457 stress. However, the convex side of bending area (the intersection between the San Andreas
458 Fault and Garlock Fault) is also at a higher stress state than those at straight segments. The
459 East California Shear Zone (ECSZ) where there are arrays of sub-parallel fault traces are in a
460 low stress state. We note that, although this is only a 2-D simplification and all faults are of
461 the same strength (viscosity), the emergent stress field has been so complex due to the
462 interaction of a fault network.

463

5 Conclusions

464 We compare the effects of both post-processing and pre-processing methods in dampening
465 down stress perturbations, which are introduced by mixed materials of strength contrast in
466 one element for the Particle-in-cell Finite Element method. The classical post-processing
467 methods alone cannot eliminate stress oscillations. Instead, with the pre-conditioning
468 methods to reduce the interface strength contrast first, the Post-local method can generate
469 a relatively accurate stress field close to analytical solutions. For the pre-conditioning methods,
470 the global method with harmonic-mean averaging method (GHM) uses the least
471 computational cost, but only helps reduce noises in models interpolated with Q_1 element
472 rather than Q_2 element. Both methods of unification of properties in one element (UnE) and
473 averaging at gauss quadrature points (AGP) need look for mixed-material cells first, but the
474 AGP method takes more time in finding and averaging particles within specified distances to
475 the gauss integration point. For the case of simple shear (fault is parallel to the velocity
476 direction), the UnE method is the best option, but for other more generic models, the AGP
477 method is the preferred. The optimal sampling distance to the gauss point is one element
478 length for both Q_2 and Q_1 elements. For models dominated by pure shear mode, the
479 arithmetic mean is better than the harmonic mean for the AGP method. Additionally, the
480 classical superconvergent point recovery (SPR) method can be further utilized to refine results
481 processed by combined pre-processing and Post-local methods, although the SPR method
482 alone cannot effectively remove stress perturbations. The synthetic model test demonstrates
483 that, with complex fault geometries, the AGP map can produce a relatively smoother stress

484 field than that in a raw result with Post-local projection.

485

486 **Acknowledgements**

487 We thank Australian Research Council for funding this research under Discovery Grant

488 DP170103350. Testing models are run with the assistance of resources from the National

489 Computational Infrastructure (NCI), as well as the Pawsey Supercomputing Centre.

490

491

492

6 References

493

494 Atwater, T., Stock, J., 1998. Pacific-North America plate tectonics of the Neogene southwestern
495 United States: an update. *International Geology Review* 40, 375-402.

496 Bird, P., 2009. Long - term fault slip rates, distributed deformation rates, and forecast of
497 seismicity in the western United States from joint fitting of community geologic, geodetic, and
498 stress direction data sets. *Journal of Geophysical Research: Solid Earth* 114.

499 Braun, J., Sambridge, M., 1994. Dynamical Lagrangian Remeshing (Dlr) - a New Algorithm for
500 Solving Large-Strain Deformation Problems and Its Application to Fault-Propagation Folding.
501 *Earth and Planetary Science Letters* 124, 211-220.

502 DeMets, C., Dixon, T.H., 1999. New kinematic models for Pacific-North America motion from 3
503 Ma to present, I: Evidence for steady motion and biases in the NUVEL-1A model. *Geophysical*
504 *Research Letters* 26, 1921-1924.

505 Deubelbeiss, Y., Kaus, B.J.P., 2008. Comparison of Eulerian and Lagrangian numerical
506 techniques for the Stokes equations in the presence of strongly varying viscosity. *Physics of*
507 *the Earth and Planetary Interiors* 171, 92-111.

508 Gizzi, F.T., 2015. Worldwide trends in research on the San Andreas Fault System. *Arabian*
509 *Journal of Geosciences* 8, 10893-10909.

510 Harlow, F.H., 1964. The particle-in-cell computing method for fluid dynamics. *Methods Comput.*
511 *Phys.* 3, 319-343.

512 Hinton, E., Campbell, J.S., 1974. Local and global smoothing of discontinuous finite element
513 functions using a least squares method. *International Journal for Numerical Methods in*
514 *Engineering* 8, 461-480.

515 Hughes, T.J.R., 2012. *The finite element method: linear static and dynamic finite element*
516 *analysis*. Courier Corporation.

517 Lenardic, A., Moresi, L.N., Mühlhaus, H., 2003. Longevity and stability of cratonic lithosphere:
518 Insights from numerical simulations of coupled mantle convection and continental tectonics.
519 *Journal of Geophysical Research: Solid Earth* 108, n/a-n/a.

520 May, D.A., Moresi, L., 2008. Preconditioned iterative methods for Stokes flow problems arising
521 in computational geodynamics. *Physics of the Earth and Planetary Interiors* 171, 33-47.

522 Meade, B.J., Hager, B.H., 2005. Block models of crustal motion in southern California
523 constrained by GPS measurements. *J Geophys Res-Sol Ea* 110.

524 Moresi, L., Dufour, F., Mühlhaus, H.B., 2003. A Lagrangian integration point finite element
525 method for large deformation modeling of viscoelastic geomaterials. *Journal of Computational*
526 *Physics* 184, 476-497.

527 Moresi, L., Quenette, S., Lemiale, V., Meriaux, C., Appelbe, B., Mühlhaus, H.B., 2007.
528 *Computational approaches to studying non-linear dynamics of the crust and mantle*. *Physics of*

529 the Earth and Planetary Interiors 163, 69-82.

530 Moresi, L., Zhong, S., Gurnis, M., 1996. The accuracy of finite element solutions of Stokes's
531 flow with strongly varying viscosity. *Physics of the Earth and Planetary Interiors* 97, 83-94.

532 Ramberg, H., 1962. Contact strain and folding instability of a multilayered body under
533 compression. *Geologische Rundschau* 51, 405-439.

534 Sadeghirad, A., Brannon, R.M., Burghardt, J., 2011. A convected particle domain interpolation
535 technique to extend applicability of the material point method for problems involving massive
536 deformations. *International Journal for Numerical Methods in Engineering* 86, 1435-1456.

537 Schmeling, H., Babeyko, A.Y., Enns, A., Faccenna, C., Funiciello, F., Gerya, T., Golabek, G.J.,
538 Grigull, S., Kaus, B.J.P., Morra, G., 2008. A benchmark comparison of spontaneous subduction
539 models—Towards a free surface. *Physics of the Earth and Planetary Interiors* 171, 198-223.

540 Sulsky, D., Chen, Z., Schreyer, H.L., 1994. A Particle Method for History-Dependent Materials.
541 *Computer Methods in Applied Mechanics and Engineering* 118, 179-196.

542 Thielmann, M., May, D.A., Kaus, B.J.P., 2014. Discretization Errors in the Hybrid Finite Element
543 Particle-in-cell Method. *Pure and Applied Geophysics* 171, 2165-2184.

544 Wallstedt, P.C., Guilkey, J.E., 2011. A weighted least squares particle - in - cell method for solid
545 mechanics. *International journal for numerical methods in engineering* 85, 1687-1704.

546 Xia, D.F., Xu, S.L., Qi, F., 1999. A proof of the arithmetic mean-geometric mean-harmonic mean

547 inequalities. RGMIA research report collection 2.

548 Zhong, S., 1996. Analytic solutions for Stokes' flow with lateral variations in viscosity.

549 Geophysical Journal International 124, 18-28.

550 Zienkiewicz, O.C., Zhu, J.Z., 1992a. The superconvergent patch recovery and a posteriori error

551 estimates. Part 1: The recovery technique. International Journal for Numerical Methods in

552 Engineering 33, 1331-1364.

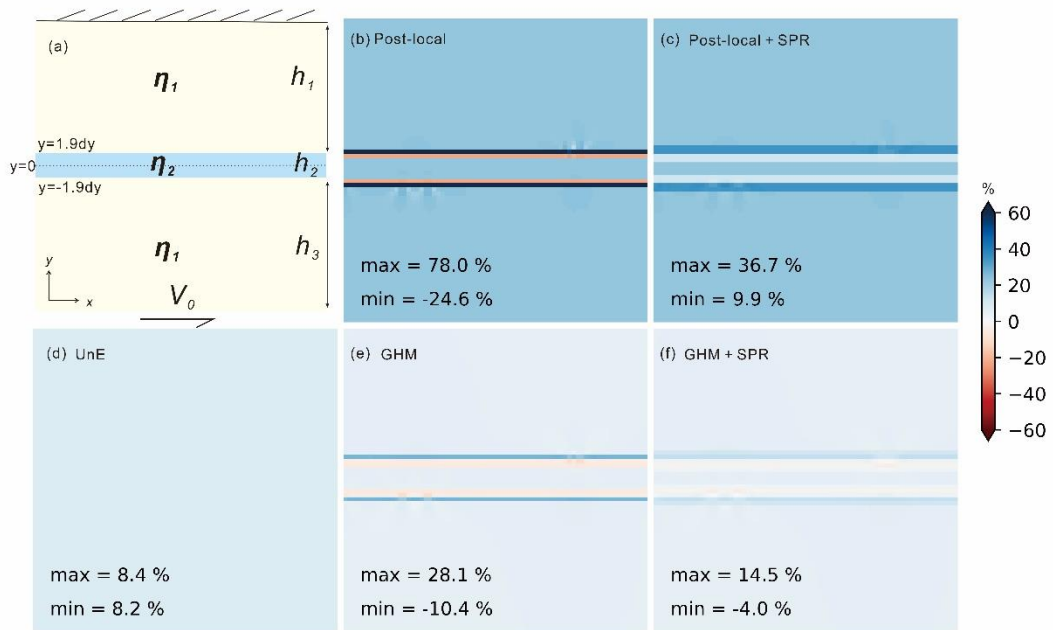
553 Zienkiewicz, O.C., Zhu, J.Z., 1992b. The superconvergent patch recovery and a posteriori error

554 estimates. Part 2: Error estimates and adaptivity. International Journal for Numerical Methods

555 in Engineering 33, 1365-1382.

556

557



559

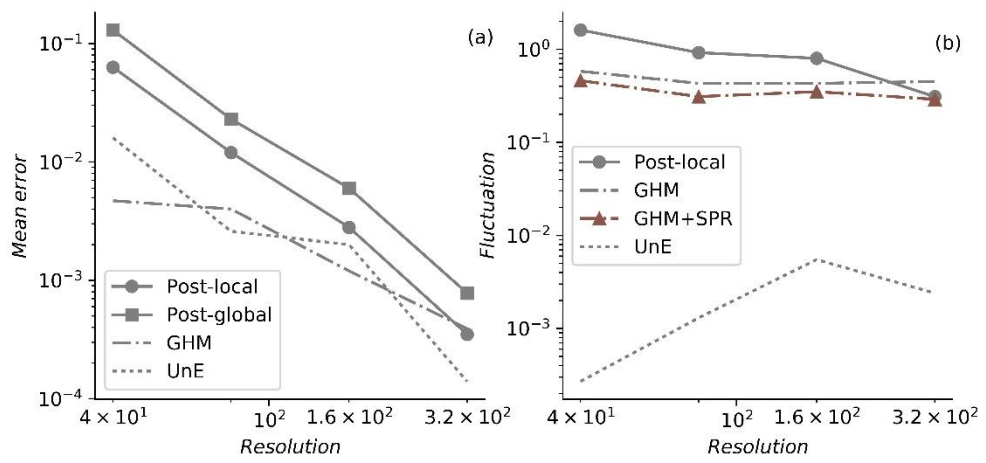
560 **Figure 1.** Simple shear model setup (a) and results of shear stress for Post-local projection

561 (b), further applying SPR method on Post-local results (c), UnE pre-processing method (d),

562 GHM pre-processing method (e), and post-processing with SPR on GHM results (f). The color

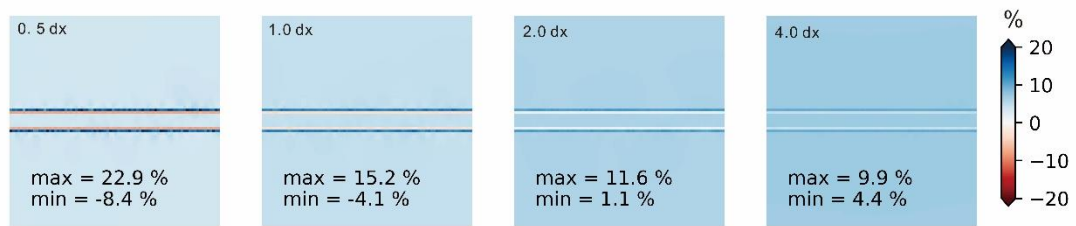
563 bar shows the normalized relative error. The mesh consists of 80×80 regular Q_1 elements.

564



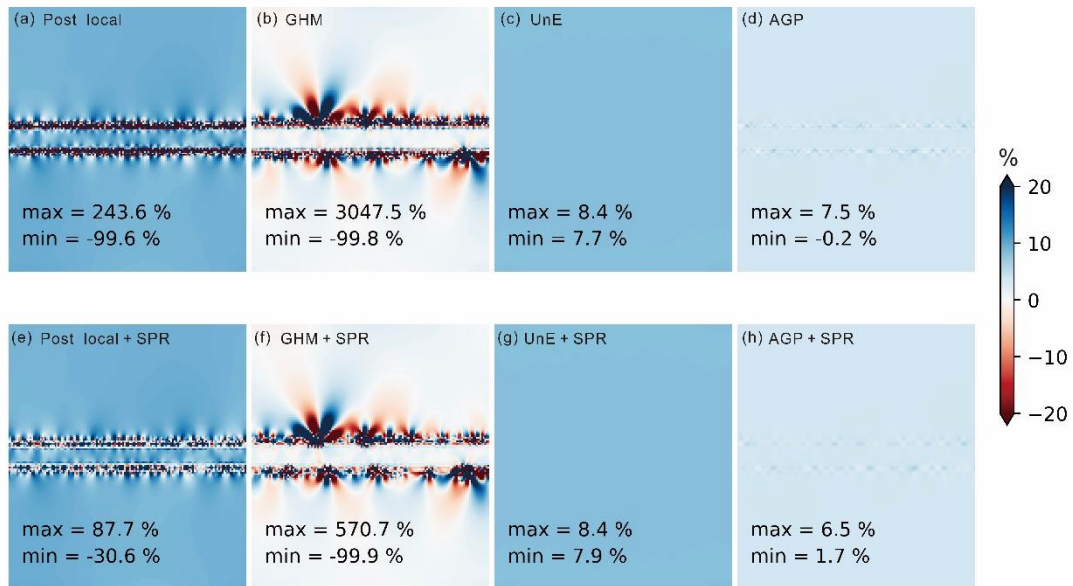
566 **Figure 2.** The mean error (a) and relative fluctuation range (b) versus resolution of element
 567 numbers.

569



570 **Figure 3.** Recovered shear stress by using the APG pre-processing method with the
571 sampling radius of 0.5 (a), 1 (b), 2 (c) and 4 (d) times of one element size. The color bar
572 shows the normalized relative error. The mesh consists of 80×80 rectangular Q_1 elements.
573 The relative error ranges decrease with sampling radius.

574



575

576 **Figure 4.** The simple shear model calculated with Q_2 element with the mesh resolution of 80

577 $\times 80$. The upper panel shows results from the Post-local projection (a), and the pre-

578 processing methods, including the GHM (b), the UnE (c), and the AGP methods (d); lower

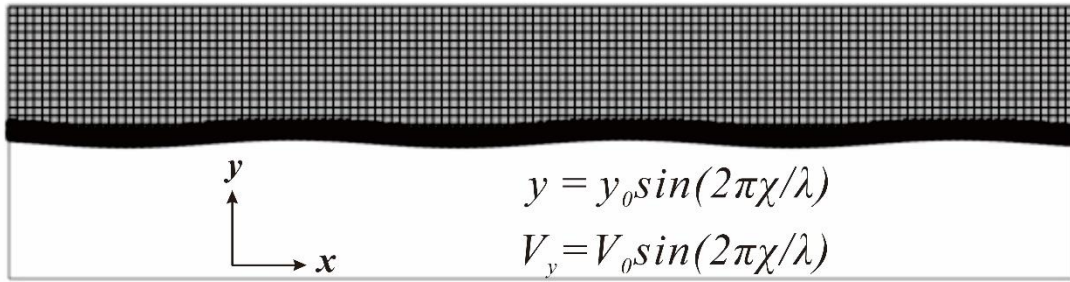
579 panel illustrates corresponding results after further processing with the SPR method. The

580 color bar shows the normalized relative error.

581

582

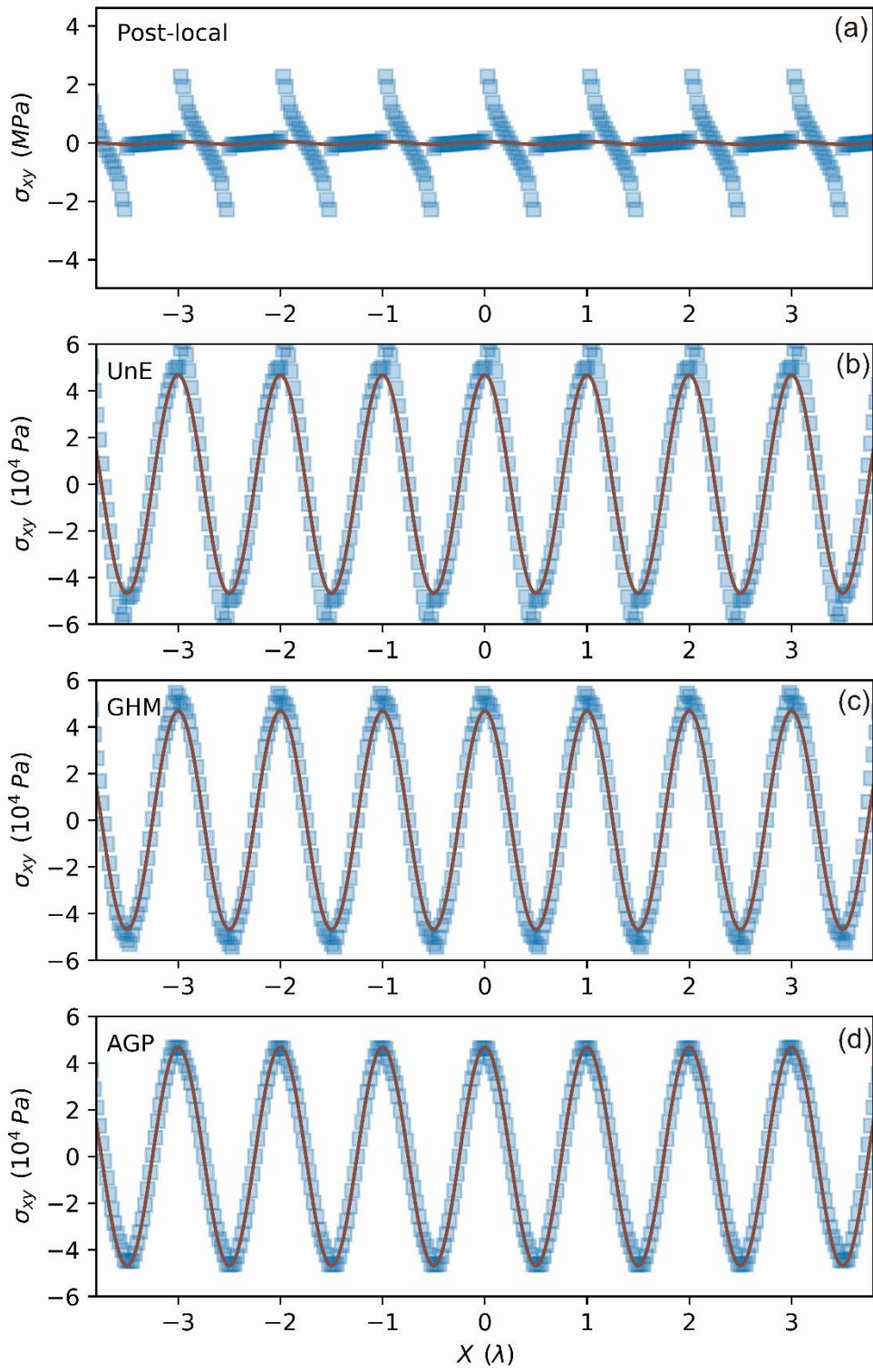
583



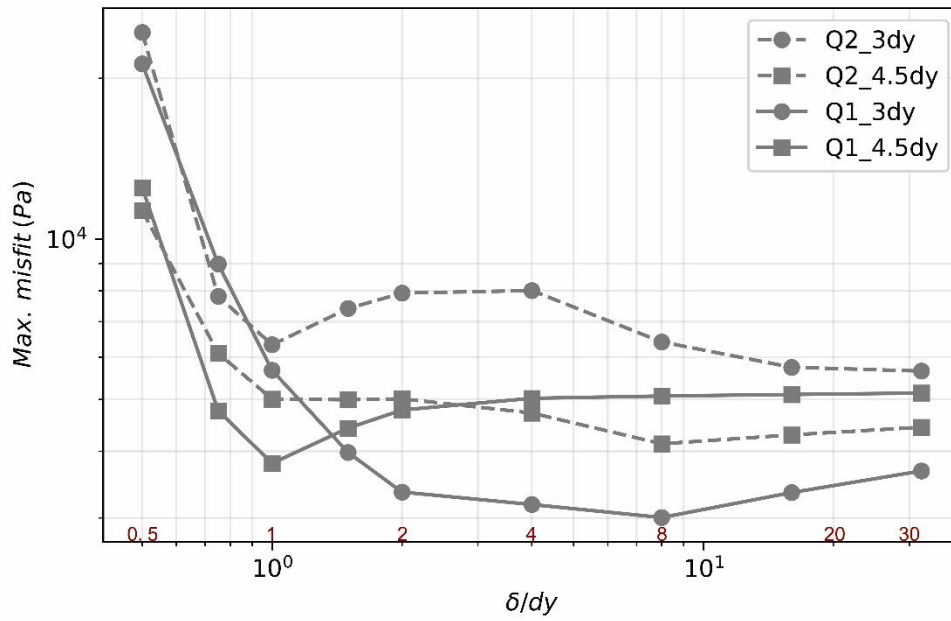
584

585 **Figure 5.** Sketch model setup that shows a thin sinusoidal competent layer (black)
586 embedded in a background incompetent rocks (grey). The rectangular mesh is overlying the
587 material field. This is only part of the model domain to show details of the mesh and mixed-
588 material elements.

589



591 **Figure 6.** Comparing analytical results for the folding model (solid line) with numerical
592 solutions (square markers) that are processed with different methods: Post-local (a), UnE (b),
593 GHM (c) and AGP(d). The sampling range in AGP is one element size.
594



595

596 **Figure 7.** The maximum misfit between analytical solution for the folding model and results

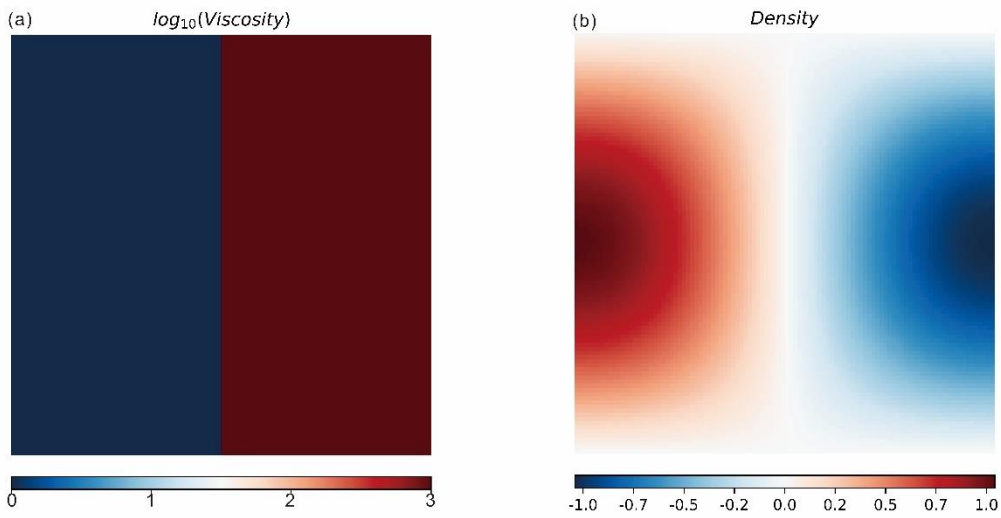
597 processed by the AGP method of different sampling radius for Q_1 (solid line) and Q_2

598 elements (dashed line). The mesh resolution for Q_1 is 320×320 , Q_2 of resolution 160×160 .

599 The observation points are at $y = 3dy$ and $4.5dy$ with the bottom of the model at $y = 0$.

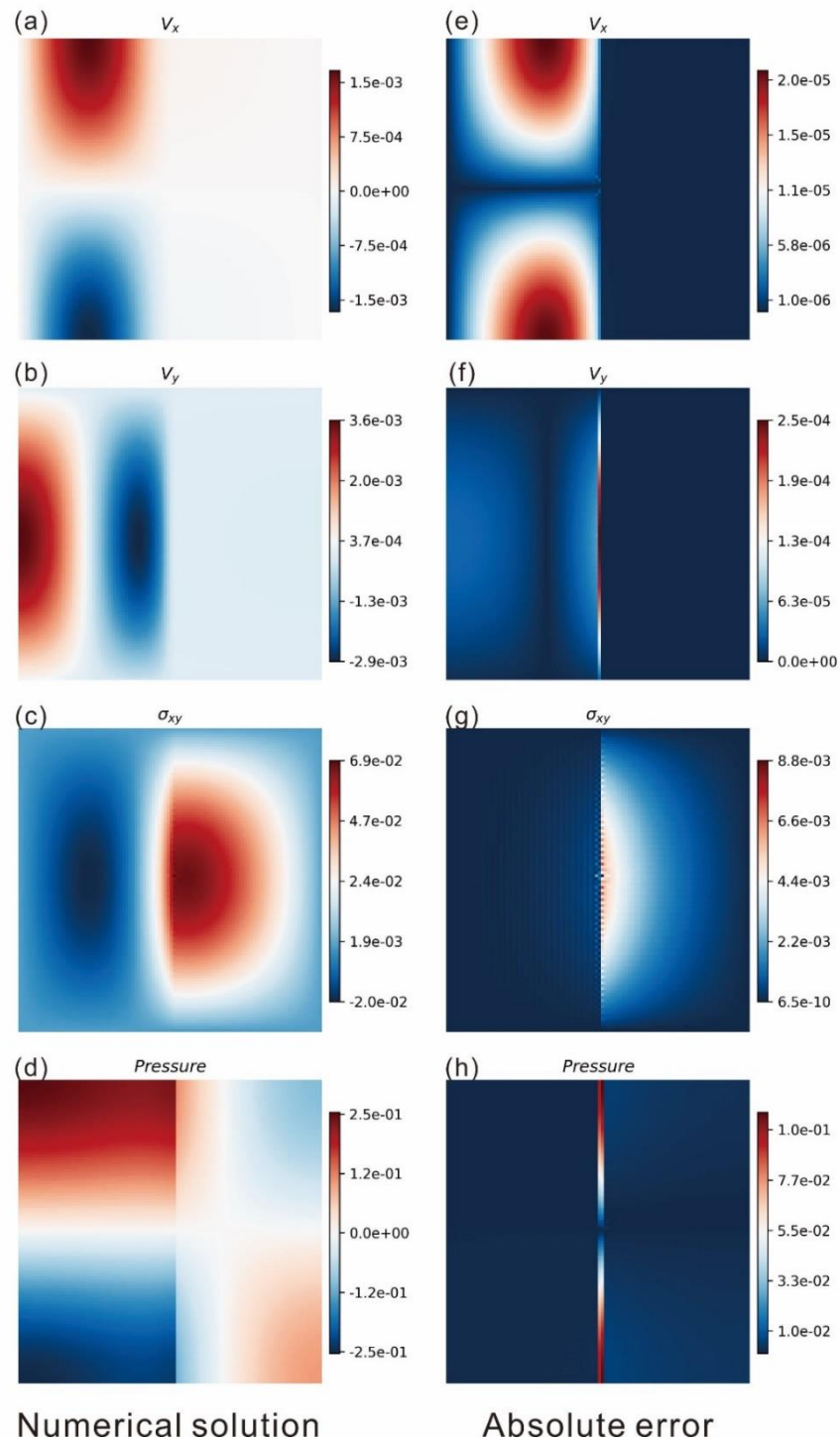
600

601



602 **Figure 8.** model setup for the SolCx benchmark.

603

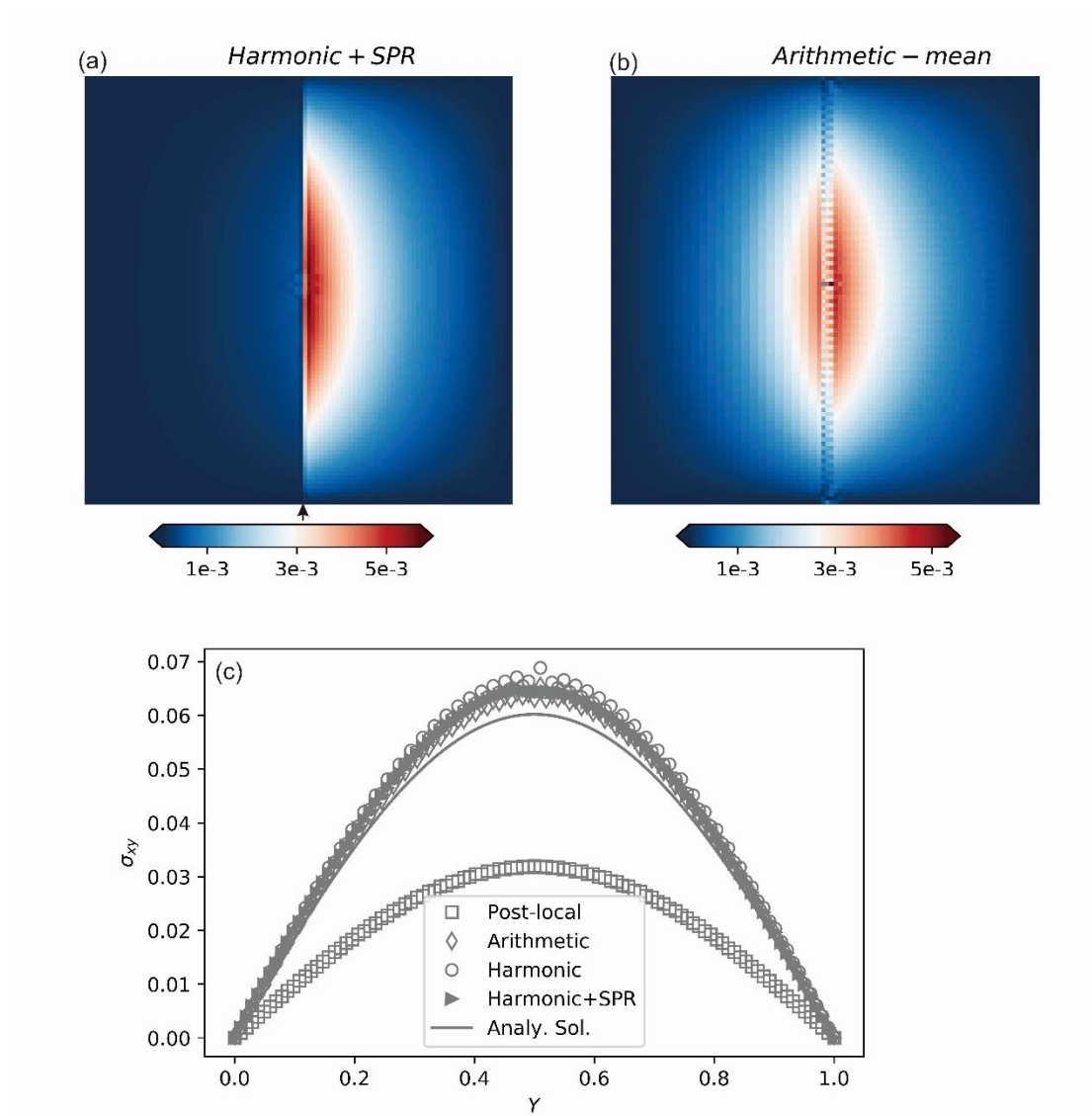


604

605 **Figure 9.** Left panel shows numerical solutions for V_x (a), V_y (b), σ_{xy} (c) and pressure (d)

606 which are processed with the AGP method (search length is one element size), and right

607 panel demonstrates corresponding absolute error.



609

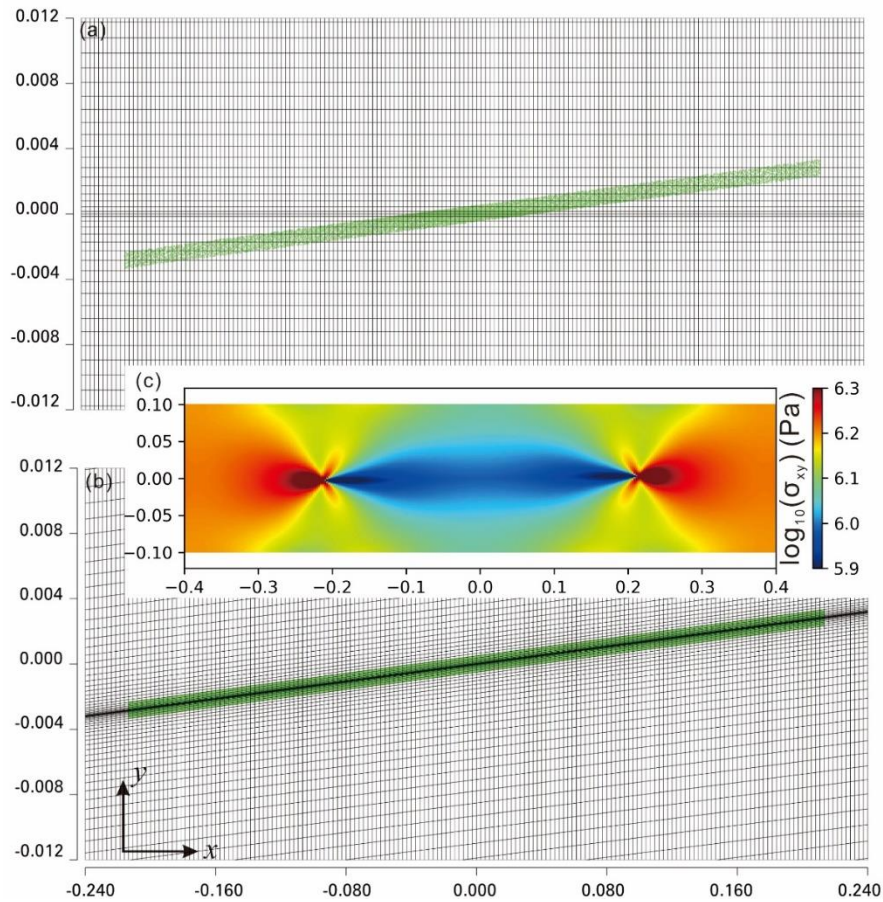
610 **Figure 10.** The absolute shear stress error for different methods: (a) applying SPR methods

611 to results obtained by the AGP method with harmonic mean, and (b) only AGP with arithmetic

612 mean. (c) The shear stress along the profile at $x=0.51$, the position of which is marked by an

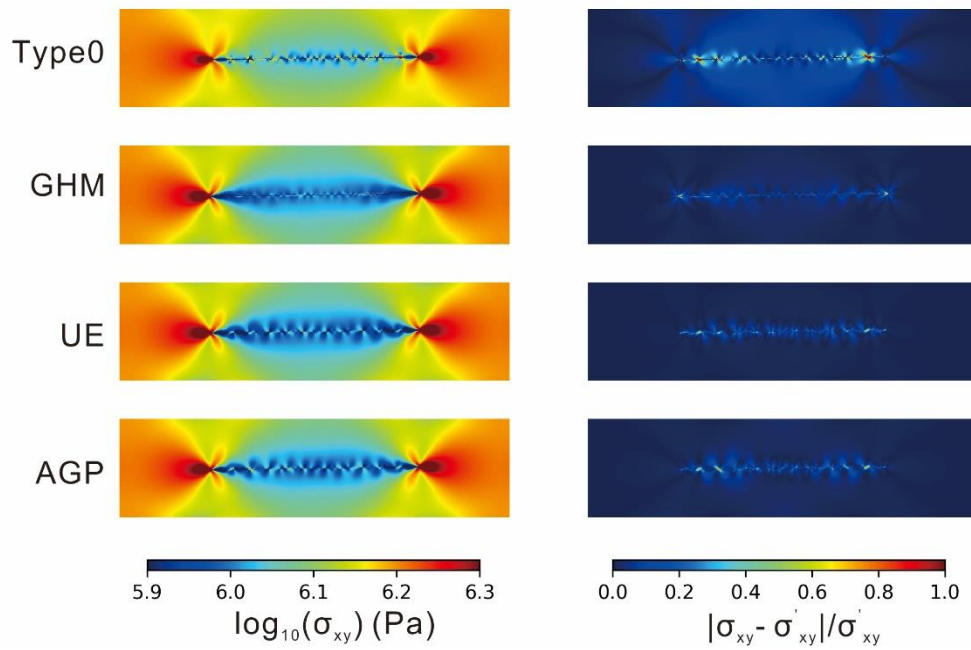
613 arrow in (a), for different methods.

614

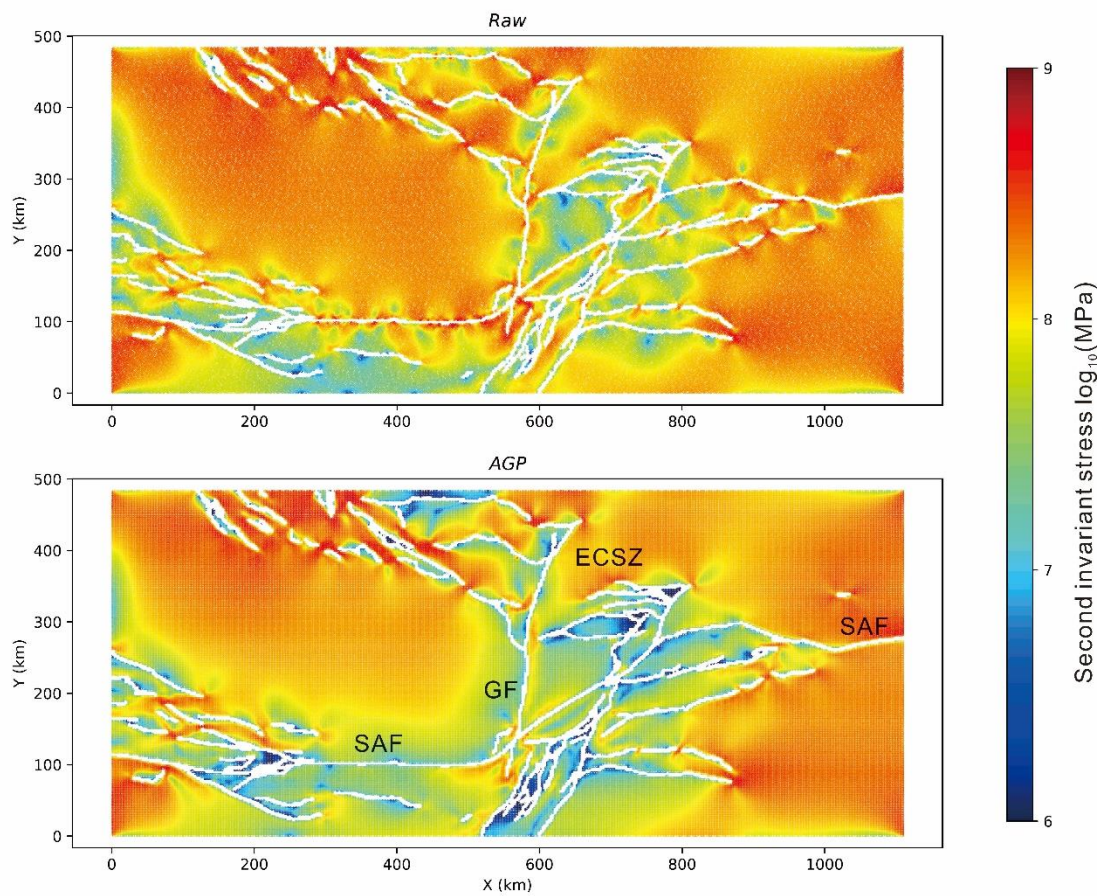


615

616 **Figure 11.** Models with an embedded weak zone (green particles at $x=-0.21-0.21$) that is at
 617 a low angle ($\sim 10^\circ$) to the x -axis. The model size is $80 \text{ km} \times 20 \text{ km}$, and the length shown
 618 here is normalized by 100 km. Q_1 elements of 300×150 is applied, and the mesh in the
 619 central part that is close to the weak zone is refined in y direction (a and b). The mesh edges
 620 in (a) is not aligned with the weak zone strike, while that in (b) is designed to be aligned
 621 with material interfaces. That means material mixing happens in (a) but not in (b). Note that
 622 the meshes in (a) and (b) are zoomed to see details of the mesh structure and are only part
 623 of the calculation domain. The whole-domain shear stress of the case in (b) is shown in (c).
 624 The results of the case in (a) is illustrated in Figure 12.



626 **Figure 12.** The results based on the mesh in Figure 11a. which has mixing materials in
 627 elements. The left panel shows absolute shear stress and the right panel shows the relative
 628 difference with respect to the reference model for different post- and pre-processing
 629 methods. σ'_{xy} is the stress in the reference model.
 630



631

632 **Figure 13.** The second invariant stress of the synthetic model of complex fault geometries
 633 that are derived from the San Andreas Fault system in California ([Bird, 2009](#)). The white belts
 634 are the mapped faulting area, and only stresses in off-fault areas are shown. Briefs for major
 635 faults: SAF- San Andreas Fault, GF-Garlock Fault, ECSZ-East California Shear zone.