# Use of machine learning to estimate statistics of the posterior distribution in probabilistic inverse problems - an application to airborne EM data.

T. M. Hansen<sup>1</sup>and C. C. Finlay<sup>2</sup>
 <sup>1</sup>Department of Geoscience, Aarhus University, Denmark
 <sup>2</sup>DTU Space, Technical University of Denmark, Kgs. Lyngby, Denmark
 Key Points:
 A machine learning approach for the probabilistic solution of inverse problems by directly estimating posterior statistics of any continuous or discrete feature of the posterior distribution.

• Allows the use of complex prior information and noise models.

1

2

3

7

8

9

10

11

Demonstrated on nonlinear probabilistic inversion of airborne electromagnetic; en ables analysis of more than 10<sup>5</sup> 1D soundings per second.

Corresponding author: Thomas Mejer Hansen, tmeha@geo.au.dk

#### 14 Abstract

The solution to a probabilistic inverse problem is the posterior probability distribution 15 for which a full analytic expression is rarely possible. Sampling methods are therefore 16 often used to generate a sample from the posterior. Decision-makers may be interested 17 in the probability of features related to model parameters (for example existence of a pol-18 lution or the cumulative clay thickness) rather than the individual realizations themselves. 19 Such features and their associated uncertainty, are simple to compute once a sample from 20 the posterior distribution has been generated. However, sampling methods are often as-21 sociated with high computational costs, especially when the prior and posterior distri-22 bution is non-trivial (non-Gaussian), and when the inverse problem is non-linear. Here 23 we demonstrate how to use a neural network to directly estimate posterior statistics of 24 continuous or discrete features of the posterior distribution. The method is illustrated 25 on a probabilistic inversion of airborne EM data from Morrill Nebraska, where the for-26 ward problem is nonlinear and the prior information is non-Gaussian. Once trained the 27 application of the network is fast, with results similar to those obtained using much slower 28 sampling methods. 29

## 30 1 Introduction

A key challenge in geoscience is that of combining different kinds of geo-information into one geo-model, typically describing the subsurface. This information can be direct information about geological processes, spatial variability, or it can be indirect information from measurements of properties related to the subsurface, such as geophysical data. Ideally, when such a geo-model has been established, one should be able to quantify information about specific features related to the geo-model, consistent with all information.

This integration of geo-information is typically solved using inverse problem the-38 ory (Tarantola & Valette, 1982a; Menke, 2012). Fast deterministic methods exist and 39 have been widely used. For such methods, the goal is to obtain one optimal model, such 40 as the simplest possible model, consistent with available information, typically in the form 41 of observed data (Tikhonov, 1963; Menke, 2012; Constable et al., 1987). In practice, in 42 part due to noise on data and model nonlinearities and imperfections, infinitely many 43 models exist that will be consistent with data, and the deterministic approach can in gen-44 eral not account properly for such uncertainty. 45

Probabilistic inversion methods can, in principle, take into account arbitrarily com-46 plex information, and integrate the information into one consistent model, as given by 47 the posterior probability distribution. A full analytic expression of the posterior distri-48 bution is rarely possible. Instead, sampling methods can be used to generate a sample 49 of the posterior, which is a collection of realizations drawn from the posterior distribu-50 tion. From such a sample, the posterior statistics of any feature related to the model pa-51 rameters can be computed. The probabilistic approach is therefore ideal for decision-makers 52 for uncertainty quantification, as it allows probabilistic analysis and risk assessment con-53 sistent with available information. 54

The main obstacle to applying the probabilistic methodology in practice is that sam-55 pling methods are computationally very demanding (Hastings, 1970; Mosegaard & Taran-56 tola, 1995). In some cases information about the posterior distribution can be used, for 57 example to construct a proposal distribution similar to the posterior distribution (Khoshkholgh 58 et al., 2022), or in the form of information about the gradient of the posterior distribu-59 tion (Fichtner et al., 2018), which can lead to more efficient sampling algorithms. Such 60 cases are however often based on rather simplistic choices of prior models. In general, 61 sampling-based methods typically require sampling or evaluation of a prior model, eval-62 uation of the physical forward response(s), and evaluation of a noise model, many times. 63

<sup>64</sup> One approach for reducing the computational requirements is to make use of fast <sup>65</sup> approximate forward modeling. This can be related to using simplified 1D forward mod-<sup>66</sup> eling as opposed to 3D forward models, or by using approximate physical models, which <sup>67</sup> leads to modeling errors that should be accounted for (Hansen et al., 2014; Madsen & <sup>68</sup> Hansen, 2018; Köpke et al., 2018).

Machine learning algorithms, which are fast to evaluate once trained, have also been used to approximate the forward modeling (Hansen & Cordua, 2017; Conway et al., 2019; Moghadas et al., 2020; Bording et al., 2021). Unsupervised machine learning methods, for example Generative adversarial neural networks (GANs), have been used more generally as a means of representing features in a prior dataset; once trained, these provide an efficient means of rapidly generating many prior realizations (Mosser et al., 2017; Laloy et al., 2018; Mosser et al., 2020).

Attempts have also been made to use machine learning methods to learn a map-76 ping from data to model that can directly solve the inverse problem. Röth and Taran-77 tola (1994) were amongst the first to solve an inverse problem in this way using a mul-78 tilayer perceptron neural network, and demonstrated an application of inversion of re-79 flection seismic data to obtain single estimates of 1D velocity profiles. Recently, several 80 authors have further explored this approach for directly solving a geophysical inverse prob-81 lem, making use of convolutional neural networks (Puzyrev & Swidinsky, 2019; Moghadas, 82 2020; Bai et al., 2020). A drawback of such methods is that, as in the deterministic so-83 lution of an inverse problem, they estimate only a single model, typically without accounting for uncertainty on geophysical data, and do not quantify the uncertainty on the pre-85 dicted model parameters. 86

An important step towards finding probabilistic solutions to inverse problems us-87 ing neural networks was made by Devilee et al. (1999) who considered training data sets 88 consisting of realizations from the prior distribution and the corresponding forward sim-89 ulated data with and without noise. They then used neural networks to learn a set of 90 statistics about each model parameter, including median and equidistant histogram es-91 timators. Meier et al. (2007) extended this work and used a mixture density network (MDN) 92 to estimate the parameters of a Gaussian mixture model representing a parametric dis-93 tribution that approximated the 1D marginal posterior distribution, and applied it to 94 the problem of estimating global crustal thickness maps, comparing to results obtained 95 using a Monte Carlo based sampling method. Several other applications of MDN to ap-96 proximate the posterior distribution, for different geophysical problems, have followed 97 (Shahraeeni & Curtis, 2011; de Wit et al., 2013; Earp & Curtis, 2020; Earp et al., 2020). 98

<sup>99</sup> Zhang and Curtis (2020a) argue that it may be problematic to apply such MDN's for higher dimensional inverse problems, and suggest to use variational inference (Blei et al., 2017) to estimate the mean and standard deviation of the (non-Gaussian) posterior distribution in an example of a seismic tomographic inverse problem. This method has been developed further for variational full waveform inversion and tomographic inversion using normalizing flows (Zhang & Curtis, 2020b; Zhao et al., 2022). In all these cases a uniform prior was assumed.

Attempts have also been made to use so-called invertible neural networks to simul-106 taneously estimate both the forward and inverse mapping between data and model pa-107 rameters (Ardizzone et al., 2018). This approach, which has recently been applied to geo-108 physical data by Zhang and Curtis (2021), allows the generation of multiple realizations 109 of the posterior distribution, from which properties of the posterior distribution can be 110 estimated, although constructing invertible neural networks involves more work than tra-111 ditional neural networks and involves compromises related to the flexibility of the net-112 work. 113

Here we present a method where the goal is not primarily to estimate the marginal 114 1D posterior distribution (as in works based on Meier et al. (2007); Shahraeeni and Cur-115 tis (2011); de Wit et al. (2013); Earp et al. (2020)). Instead, we propose and demonstrate 116 a machine learning-based method that provides direct estimates of any desired statis-117 tical property (continuous or discrete) of the posterior distribution, including any fea-118 ture or property that can be computed from realizations of an, in principle, arbitrarily 119 complex, prior model. This is achieved without generating realizations of the posterior 120 distribution. 121

122 Following Devilee et al. (1999) and Meier et al. (2007) we construct a finite size training data set, representing the information available in the probabilistic formulation of 123 the inverse problem, namely prior information and information about the forward model 124 and the noise. This is then used to train a neural network whose output parameterizes 125 any desired statistical property of the posterior distribution for which a log-likelihood 126 can be computed. These properties can for example represent a Gaussian, generalized 127 Gaussian, log-normal, or a mixture model distribution, representing continuous model 128 parameters. The output can also refer to the posterior probability of defined classes of 129 model features of discrete model parameters. The neural network is designed to ensure 130 that the estimated statistical properties of the posterior are similar to the same statis-131 tics derived from a sample of the posterior. Given a suitable training set the method pro-132 vides accurate information regarding properties of the posterior distribution of interest 133 in a given problem at a fraction of the computational cost of traditional sampling-based 134 approaches. 135

In the following the method is first presented for probabilistic inverse problems in 136 general; this can be considered a generalization of the ideas proposed by Devilee et al. 137 (1999) and followed up by e.g. Meier et al. (2007); Earp et al. (2020). Next, we demon-138 strate the method, applying it to non-linear probabilistic inversion of airborne electro-139 magnetic data using non-Gaussian prior models of varying complexity. We show the neu-140 ral network approach can be used to accurately estimate statistical properties of the pos-141 terior, related to both discrete and continuous model parameters, using regression and 142 classification networks. The results are compared to results obtained by calculating the 143 same statistical properties from a sample of the posterior obtained using the extended 144 rejection sampler (Hansen, 2021). 145

# 146 2 Method

Let  $\mathbf{m} = [m_1, m_2, \dots, m_{N_M}]$  represent  $N_M$  model parameters that define some properties of a system, such as for example physical properties of a geo-model.  $\mathbf{m}$  is typically represented on a grid in a Cartesian or spherical coordinate system. For example,  $\mathbf{m}$  might represent geophysical properties such as resistivity, velocity, or any other geological/geophysical/geochemical parameter.

A key issue in geosciences is how to infer information about **m** from different types 152 of available information, such as geological expert knowledge, geophysical data, well log 153 data, etc. This is generally referred to as an inverse problem. Tarantola and Valette (1982b) 154 describe the inverse problem as a problem of probabilistic conjunction of information. 155 Available information about  $\mathbf{m}$  is described in the form of probability densities and then 156 combined using conjunction of information to obtain a single probability density that 157 describes the combined information. For example, consider a case when a specific type 158 of information about structural properties is quantified by  $\rho(\mathbf{m})$ , and that information 159 from observed electromagnetic (EM) data and well logs is quantified through  $L(\mathbf{m})$ . Then 160 the conjunction of this information is given by the posterior probability distribution  $\sigma(\mathbf{m})$ . 161 which, under the assumption that the individual types of information have been obtained 162 independently, is given by 163

$$\sigma(\mathbf{m}) \propto \rho(\mathbf{m}) \cdot L(\mathbf{m}). \tag{1}$$

164	In other words, the conjunction of the independent information is proportional to
165	the product of probability densities describing each independent set of information. The
166	likelihood $L(\mathbf{m})$ is a measure of how well the data <b>d</b> computed from a specific model matches
167	observed data $\mathbf{d}_{obs}$ given noise with a specified probability distribution. Noise-free data
168	can be computed by evaluating the forward model

$$\mathbf{d} = g(\mathbf{m}),\tag{2}$$

where g is a non-linear operator that maps the model parameters into data. g typically refers to some numerical algorithm solving some physical equations (such as Maxwell's equations).

The probabilistic inverse problem is then to infer information about  $\sigma(\mathbf{m})$ , which contains the combined information of, for example, both structural prior information, through the prior  $\rho(\mathbf{m})$ , and information from observed geophysical data, through  $L(\mathbf{m})$ .

A general approach (that allows using a non-linear forward model and non-Gaussian prior) for solving probabilistic formulated inverse problems is use of sampling methods to sample the posterior distribution, Eqn. 1, (Metropolis et al., 1953; Hastings, 1970; Geman & Geman, 1984; Green, 1995; Mosegaard & Tarantola, 1995; Laloy & Vrugt, 2012; Hansen et al., 2013, 2016). Unfortunately, such sampling methods can be extremely computationally demanding, to the point where they cannot be practically applied. They rely on solving the forward problem, Eqn. 2, many (often millions of) times.

Some algorithms make implicit assumptions about the prior model, such as a layered subsurface (Malinverno, 2002; Sambridge et al., 2013), while others, such as the classical rejection sampler and Metropolis algorithm (Hastings, 1970) require that both the prior and likelihood can be evaluated. This typically leads to using relatively simple prior models.

The extended variations of the Metropolis algorithm (Mosegaard & Tarantola, 1995) and the rejection sampler (Hansen et al., 2016; Hansen, 2021) do not require that an analytical description of the prior exists, as evaluation of the prior is not needed. It is sufficient that an algorithm exists that can generate a realization from the prior. This opens up the possibility of using a variety of more complex prior models, based on for example geostatistical simulation-based methods (Hansen et al., 2008, 2012).

193

# 2.1 Properties related to geophysical model parameters.

The model parameters **m** typically refer to physical parameters (e.g. resistivity when dealing with EM data, or elastic properties when dealing with seismic data). In practice, decision makers may be more interested in related features, or specific questions, such as "What is the chance of penetrating a specific lithology when drilling'?" (Scales & Snieder, 1997). Such features or occurences of events will be referred to through **n**.

In general the relation between **m** and **n** can be complex and is formally described by a joint prior distribution  $\rho(\mathbf{m}, \mathbf{n})$ . This can for example be the case if **n** refers to subsurface lithology, and **m** to a geophysical property. This has been widely studied in the inversion of reflection seismic data, where information about geophysical properties is often assumed dependent on lithology, such that  $\rho(\mathbf{m}, \mathbf{n}) = \rho(\mathbf{n})\rho(\mathbf{m}|\mathbf{n})$  (Bosch et al., 2010; Grana & Della Rossa, 2010; Rimstad et al., 2012). A more general formulation of Eqn. 1, describing information on both **m** and **n** is then

$$\sigma(\mathbf{m}, \mathbf{n}) \propto \rho(\mathbf{m}, \mathbf{n}) \cdot L(\mathbf{m}, \mathbf{n}), \tag{3}$$

given the available joint prior information, the forward model, and the noise. The cor responding forward problem, generalizing Eqn. 2, takes the form

$$\mathbf{d} = g(\mathbf{m}, \mathbf{n}). \tag{4}$$

Sometimes the relation between  $\mathbf{m}$  and  $\mathbf{n}$  is so simple that  $\mathbf{n}$  can be computed from 208 **m** through a mapping function  $\mathbf{n} = h(\mathbf{m})$ . For example, **n** can refer to the volume of 209 a reservoir (a scalar) obtained from a high dimensional set of geophysical model param-210 eters  $\mathbf{m}$ . Or,  $\mathbf{n}$  can refer to the cumulative thickness of layers with a resistivity ( $\mathbf{m}$ ) above 211 some threshold. Another example is when  $\mathbf{m}$  refers to properties of a groundwater model. 212 Then flow modeling based on a set of realizations from the posterior, can be used to prop-213 agate uncertainties into for example, the arrival time of polluted groundwater  $(\mathbf{n})$  at a 214 specific location (Vilhelmsen et al., 2019). Such a focus on related features and proper-215 ties derived from the posterior distribution, rather than the posterior distribution over 216 the geophysical parameter  $\sigma(\mathbf{m})$  itself, is discussed by Scheidt et al. (2015). 217

The sampling algorithms described above can be used to generate a sample from  $\sigma(\mathbf{m}, \mathbf{n})$  from which statistical analysis of any feature related to  $\sigma(\mathbf{m}, \mathbf{n})$  can be computed.

The goal here is however not to generate realizations of the posterior distribution, 220 but instead to compute directly statistical properties of the posterior distribution sim-221 ilar to those that would be obtained by computing it directly from a sample of the posterior distribution. In other words, given a sample  $\hat{\mathbf{n}}$  of the posterior,  $\sigma(\mathbf{n})$ , the goal is 223 to compute parameters  $\Theta$  that define a desired statistical property of  $\sigma(\mathbf{n})$ . For exam-224 ple, if **n** refers to a discrete parameter with  $N_o$  possible outcomes, then  $\Theta = [\theta_1, ..., \theta_{N_o}]$ 225 could refer to the probability of realizing each possible outcome. If  $\mathbf{n}$  refers to a contin-226 uous parameter,  $\Theta = [\theta_0, \mathbf{C}_{\theta}]$  could represent the mean and covariance of a multivari-227 ate Gaussian distribution.  $\Theta = [\theta_0, \theta_1, \theta_2]$  could represent the mean, variance and power 228 of a generalized 1D Gaussian distribution.  $\Theta = [\theta_0]$  could represent the rate of a Pois-229 son distribution.  $\Theta = [\theta_0, \theta_1]$  could represent a Binomial distribution. 230

Assume that a sample  $\hat{\mathbf{n}}$  of  $\sigma(\mathbf{n})$  is available. The optimal values of  $\Theta$  can be found maximizing the likelihood,  $L_{\Theta}$ , that each realization of the posterior,  $\hat{\mathbf{n}}^{i*}$ , is a realization of the probability distribution (described by the parameter(s)  $\Theta$ )  $f(\hat{\mathbf{n}}^{i*}|\Theta)$ , given as

$$L_{\Theta} = f(\hat{\mathbf{n}}|\Theta) = \prod_{i=1}^{N_{\sigma}} f(\hat{\mathbf{n}}^{i*}|\Theta),$$
(5)

where  $N_{\sigma}$  is the number of independent realizations of  $\hat{\mathbf{n}}$ . The specific choice of  $f(\hat{\mathbf{n}}^{i*}|\Theta)$ depends on the type of statistical parameters to be estimated. Examples will be given below. Maximization of Eqn. 5 is equivalent to minimizing the negative log-likelihood (which we refer to as the loss,  $J_{\Theta}$ ):

$$J_{\Theta} = -\log(\prod_{i=1}^{N_{\sigma}} f(\hat{\mathbf{n}}^{i*}|\Theta)) \tag{6}$$

$$= -\sum_{i=1}^{N_{\sigma}} \log(f(\hat{\mathbf{n}}^{i*}|\Theta)).$$
(7)

<sup>239</sup> Minimization of the loss function, Eqn. 7, can be used to obtain estimates of the <sup>240</sup> parameters  $\Theta$  representing statistical properties of  $\sigma(\mathbf{n})$ .

Here a method is proposed that allows direct computation of the parameters,  $\Theta$ , 241 that describe statistical properties of  $\sigma(\mathbf{m}, \mathbf{n})$ , using a neural network trained on a data 242 set containing a sample of the known information (including the prior, forward, noise and 243 modeling errors), without ever generating realizations from  $\sigma(\mathbf{m}, \mathbf{n})$ . The approach follows the basic strategy proposed by Devilee et al. (1999), and consists of two steps: A) 245 construction of a training data set, and B) construction and training of a neural network. 246 This is done once. Then, the trained machine learning algorithm can be applied, very 247 efficiently to compute desired properties of the posterior distribution, for potentially many 248 sets of observed data. 249

#### 250 2.2 A: Construction of training data set

Eqn. 4 describes the forward problem of computing noise free data. The forward problem describing simulation of data including noise,  $\mathbf{d}_{sim}$ , is

$$\mathbf{d}_{sim} = g(\mathbf{m}, \mathbf{n}) + r(\mathbf{m}, \mathbf{n}) = \mathbf{d} + r(\mathbf{m}, \mathbf{n}), \tag{8}$$

where  $r(\mathbf{m}, \mathbf{n})$  represent noise that can be related to the model and features. Often geophysical data **d** depends only directly on the physical parameters, in which case  $g(\mathbf{m}, \mathbf{n}) = g(\mathbf{m})$ .

Let  $\mathbf{M}^* = [\mathbf{m}^{1*}, \mathbf{m}^{2*}, ..., \mathbf{m}^{N_T*}]$  and  $\mathbf{N}^* = [\mathbf{n}^{1*}, \mathbf{n}^{2*}, ..., \mathbf{n}^{N_T*}]$  represent  $N_T$  realizations of  $\rho(\mathbf{m}, \mathbf{n})$ . Let  $\mathbf{D}^* = [\mathbf{d}^{1*}, \mathbf{d}^{2*}, ..., \mathbf{d}^{N_T*}]$  represent the corresponding  $N_T$  noise free data, obtained by evaluating Eqn. 4. Finally let  $\mathbf{D}_{sim}^* = [\mathbf{d}_{sim}^{1*}, \mathbf{d}_{sim}^{2*}, ..., \mathbf{d}_{sim}^{N_T*}]$  represent  $N_T$  corresponding realizations of simulated noisy data, following Eqn. 8. This constitutes a training data set

$$\mathbf{T} = [\mathbf{N}^*; \mathbf{M}^*; \mathbf{D}^*; \mathbf{D}^*_{sim}], \tag{9}$$

that can be obtained by 1) sampling the prior, 2) solving the forward problem, 3) simulation of the noise.

The sample **T** in Eqn. 9 represents the available information (prior, physics of the forward model, noise) in so far as it can be represented by a finite sample of size  $N_T$ . The larger the sample, the more complete the representation of the available information.

266

275

# 2.3 B: Construct and train a neural network to estimate relevant statistics of $\sigma(m, n)$

The goal is to design and train a neural network to estimate  $\Theta$  directly from realizations of simulated data including noise  $\mathbf{d}_{sim}^{i*}$ . In principle any machine learning method capable of regression and/or classification, such as regression trees and support vector machines (Bishop et al., 1995), can be used to estimate the mapping  $\mathbf{d}_{sim}^* \mapsto \Theta$  which after training can be used on real data to evaluate  $\mathbf{d}_{obs} \mapsto \Theta$ . Here we use make use of a fully connected artificial neural network. The presented approach builds on earlier work by Röth and Tarantola (1994), Devilee et al. (1999) and Meier et al. (2007).

#### 2.3.1 The structure of the neural network

A neural network can be described in terms of an input layer, an inner part of the neural network (which can consist of many layers, referred to as hidden layers), and an output layer. The input layer here represents the training data, which include noise, and consists of  $N_d$  neurons. The output layer has  $N_{\theta}$  neurons representing the statistical parameters describing a distribution characterizing the features or properties of the posterior distribution that one wishes to predict.

The inner part of the network can be either simple or complex, and it can consist of either (fully) connected layers of neurons, convolutional layers, or combinations of these and other types of layers depending on the application. Here a fully connected neural network is considered as it has been demonstrated that such a neural network, with at least one hidden layer, can approximate any continuous function with arbitrary accuracy, when the number of hidden units is large enough (Hornik et al., 1990).

Each neuron has a number of adjustable parameters, the weights  $w_i$  (one for each neuron in the previous layer), and a bias b, as well as an activation function  $\Psi$ . All neurons in one layer are fully connected to all neurons in the following layer. The input for a neuron (except for the first layer where the input is  $\mathbf{d}_{sim}^*$ ) is the output of the neurons in the previous layer, and the output y of a neuron in response to inputs  $x_i$ , is given by

$$y = \Psi\left(\sum_{i} (w_i * x_i) + b\right).$$
(10)

For a specific network, with specified values for the weights and biases, one can compute the output, given some input, simply by evaluating the neurons layer by layer, starting from the input layer. See e.g. Bishop et al. (1995) for more details.

297 2.3.2 The loss function

When a neural network is trained using the training data set, its free parameters (the weight and bias of each node for a fully connected network) are adjusted to minimize a specific loss function. In the present case, the training data set consists of (when properties of  $\sigma(\mathbf{n})$  are of interest)  $\mathbf{T} = [\mathbf{N}^*; \mathbf{D}^*_{sim}]$ . The goal is to estimate  $\mathbf{d}^*_{sim} \mapsto \Theta$ rather than simply  $\mathbf{d}^*_{sim} \mapsto \mathbf{n}$ .

This is achieved by constructing a loss function with unknown parameters  $\Theta$  that describe statistical properties of the desired probability distribution, Eqn. 5, and whose parameters can be found by minimizing the loss function, Eqn. 7. The key here is to choose a loss function that is the negative log-likelihood of the property of interest as described by the parameters  $\Theta$  one wishes to estimate.

At each iteration of training the neural network, the loss is computed by applying the following steps for each dataset  $T^i = [\mathbf{n}^{i*}, \mathbf{d}_{sim}^{i*}]$  in the training data set **T**:

1. Evaluate the network using  $\mathbf{d}_{sim}^{i*}$  as input. This provides as output an estimate  $\hat{\Theta}_i$ 

- 2. Evaluate the corresponding loss,  $J^i$ , as  $J^i = -\log(f(\mathbf{n}^{i*}|\hat{\Theta}_i))$ .
- <sup>313</sup> The total loss is then given by

312

$$\mathbf{J} = \sum_{i=1}^{N_T} J^i.$$
(11)

By construction, as  $\mathbf{d}_{sim}^{i*}$  has been computed from  $\mathbf{n}^{i*}$  using Eqn. 8,  $\mathbf{n}^{i*}$  can be consid-

ered a realization of  $\sigma(\mathbf{n})$ , given the data  $\mathbf{d}_{sim}^{i*}$ , and therefore, minimizing the loss in Eqn.

<sup>316</sup> 11 leads to estimates of statistical parameters  $\Theta$  that describe  $\sigma(\mathbf{n})$ , in the same manner as would minimizing Eqn. 7 given a sample,  $\hat{\mathbf{n}}$ , from  $\sigma(\mathbf{n})$ . The difference is that the proposed method achieves this without the need to first realize the sample  $\hat{\mathbf{n}}$ .

Minimizing the loss function thus maximizes the probability that each  $\mathbf{n}^{i*}$  can be seen as a realization of the probability distribution whose parameters  $\Theta_i$  are the result of evaluating the neural network  $\mathbf{d}_{sim}^{i*} \mapsto \Theta_i$ . Note that it is crucial that data with noise  $\mathbf{d}_{sim}^{i*}$  is used for training, as opposed to using noise free data  $\mathbf{d}^{i*}$ , as this would imply ignoring noise completely, which would lead to overfitting.

In general, **n** (and/or **m**) can refer to a continuous parameter (such as velocity, resistivity, temperature, or related properties) or a discrete parameter (such as lithology type and event type). Continuous model parameters lead to a regression type problem, whereas discrete model parameters lead to a classification problem.

228 2.3.2.1 Continuous model parameters - regression We first consider the case when 229 **n** represents continuous parameters. Say we wish to estimate the mean and covariance, 330  $\hat{\mathbf{C}}_{\theta}$ , of the posterior distribution  $\sigma(\mathbf{n})$  given a set of observed data  $\mathbf{d}_{obs}$ . Assume 331 a neural network exists that outputs a set of parameters describing  $\Theta = [\hat{\Theta}_0^i, \hat{\mathbf{C}}_{\theta}^i]$ , given 332 the input  $\mathbf{d}_{sim}^i$ . The likelihood that a set of parameters from the training dataset  $\mathbf{n}^{i*}$ 333 is a realization from the multivariate Gaussian distribution  $\mathcal{N}(\hat{\Theta}_0^i, \hat{\mathbf{C}}_{\theta})$  as obtained from 334 evaluating the neural network using  $\mathbf{d}_{sim}^{i*}$  as input, is given by

$$f(\mathbf{n}^{i*}|\hat{\Theta_0}^{i}, \hat{\mathbf{C}^{i}}_{\theta}) = k_C \exp(-0.5 (\mathbf{n}^{i*} - \hat{\Theta_0}^{i})^T \hat{\mathbf{C}}_{\theta}^{i^{-1}} (\mathbf{n}^{i*} - \hat{\Theta_0}^{i})), \quad (12)$$

where  $k_C = ((2\pi)^{N_d} |\hat{\mathbf{C}}_{\theta}^i|)^{-\frac{1}{2}}$  is a normalization factor. The corresponding loss function  $J^i$  is

$$J^{i} = -\log(f(\mathbf{n}^{i*}|\hat{\Theta_{0}}^{i}, \hat{\mathbf{C}}^{i}_{\theta}))$$
(13)

$$= 0.5 (\mathbf{n}^{i*} - \hat{\Theta_0}^{i})^T \hat{\mathbf{C}_{\theta}^{i}}^{-1} (\mathbf{n}^{i*} - \hat{\Theta_0}^{i}) - \log(k_C)$$
(14)

The total loss is then given by Eqn. 11. Any neural network that minimizes this loss function, will lead to an estimate of the parameters of interest, here  $\Theta = [\hat{\Theta}_0, \hat{\mathbf{C}}_{\theta}]$ , that are computed directly without ever computing realizations of  $\sigma(\mathbf{n})$ .

To represent the posterior mean and full covariance, given  $N_m$  model parameters, 340 an output layer of  $N_{\Theta} = N_m + N_m^2$  nodes must be used. If only the posterior mean 341 and variance are estimated, an output layer of  $N_{\Theta} = N_m + N_m$  nodes is needed. If only 342 the posterior mean is of interest an output layer of  $N_{\Theta} = N_m$  nodes is needed and min-343 imizing Eqn. 14 is then similar to minimizing the widely used mean squared error loss 344 function (Bishop et al., 1995), as utilized for example in e.g. Röth and Tarantola (1994). 345 Recall, that the above scheme does not impose any assumptions on either the prior or 346 the posterior distribution which may be complex. The estimated mean and covariance 347 are simply statistical parameters of the posterior distribution, that may or may not be 348 useful for a specific use case. The quality of the obtained estimate naturally depends on 349 the complexity of the machine learning model used, and the size of the training data set, 350 which will be considered in more detail in the application presented below. 351

Other statistical parameters of the posterior can be estimated by minimizing the appropriate negative log-likelihood function for the considered probability distribution. For example, a 1D generalized probability distribution is defined by three parameters  $\Theta = [\theta_1, \theta_2, \theta_3]$ , and its probability distribution given by (Tarantola, 2005)

$$f(n^{i}|\Theta) = \frac{1}{2\theta_{2}\Gamma(1+1/\theta_{3})} \exp\left(-\left(\frac{|n^{i}-\theta_{1}|}{\theta_{2}}\right)^{\theta_{3}}\right).$$
(15)

A 1D Gaussian mixture model based on a mixture of Nc 1D Gaussian distribution, as considered by e.g. Meier et al. (2007), is defined by  $\Theta = [\mathbf{t}_1, \mathbf{t}_2, \mathbf{t}_3] = [t_1^1, ..., t_1^{Nc}, t_2^1, ..., t_2^{Nc}, t_3^1, ..., t_3^{Nc},],$ where  $\mathbf{t}_1$  refers to the mean,  $\mathbf{t}_2$  refers to the standard deviation of Nc Gaussian distribution, each with weight  $\mathbf{t}_3$ , and its probability distribution given by

$$f(n^{i}|\mathbf{t}_{1},\mathbf{t}_{2},\mathbf{t}_{3}) = \sum_{i=1}^{Nc} t_{3}^{i} \left(t_{2}^{i}\sqrt{2\pi}\right)^{-1} \exp\left(-0.5\left(\frac{n^{i}-t_{1}^{i}}{t_{2}^{i}}\right)^{2}\right).$$
 (16)

The corresponding negative log-likelihood for Eqns. 15 and 16 can trivially be obtained and used as a loss function in a neural network to estimate  $\Theta$ . In principle, any statistical parameter with a corresponding negative log-likelihood that can be computed, and used as a loss function, can be estimated using the proposed methodology.

2.3.2.2 Discrete model parameters - classification Say  $n_i$  represents a discrete parameter with  $N_o$  possible outcomes (classes). One's aim is then to estimate the posterior probability of each of the  $N_o$  classes given some data  $\mathbf{d}_{obs}$ .

Let  $\theta_i^* = [p_i^{1*}, p_i^{2*}, ..., p_i^{N_o*}]$  represent the true probabilities of  $n_i^*$  belonging to a specific class. In practice the true probability of one (the correct) class will be one, and the others zero. Further  $\hat{\theta}_i = [\hat{p}_i^{1}, \hat{p}_i^{2}, ..., \hat{p}_i^{N_o}]$  represent the corresponding predictions by the neural network of the probabilities of each class for a specific model parameter,  $n_i$ .

The likelihood of observing  $\theta_i$  given  $\hat{\theta}_i$  is then

$$f(\theta_i|\hat{\theta}_i) = \prod_{j=1}^{N_o} (\hat{p}_i^{\ j})^{p_i^{\ j*}}.$$
(17)

373

372

The corresponding loss function  $J^i$  is then

$$J^{i} = -\log(f(\theta_{i}|\hat{\theta}_{i})) = -\sum_{j=1}^{N_{o}} p_{i}^{j*} \log(\hat{p}_{i}^{j}).$$
(18)

The choice of class probabilities  $\hat{\theta}_i$  that maximizes Eqn. 17 can be found by min-374 imizing the negative log-likelihood given by the loss function, Eqn. 18, which is equiv-375 alent to the categorical cross-entropy between the two probability distributions (Bishop 376 et al., 1995). Usually, the softmax activation is used for multi-class classification prob-377 lems (while the sigmoid activation function is used for binary classification problems), 378 as it forces all probabilities to be in the range 0 to 1, and ensures that  $\sum_{j=1}^{N_o} \hat{p}_i^j = 1$ , 379 such that the output parameters can be interpreted as a probability. A neural network that estimates the mapping  $\mathbf{d}_{sim}^i \mapsto \hat{\theta}^i$  by minimizing Eqn. 18, using the softmax activation function in the output layer, therefore locates the maximum-likelihood of Eqn. 380 381 382 17, which directly estimates  $\sigma(p_i^*)$ , i.e. the posterior class probability for a discrete model 383 parameter. 384

To summarize, our proposed method involves first constructing a training data set 385 (Eqn. 9) that represents (within the limits of the size of the training data set used) the 386 known information (the prior, the forward, and the noise model), and specifically con-387 388 tains prior knowledge regarding any feature  $\mathbf{n}$ , that may be directly or indirectly related to the model parameters **m**, about which one wishes to infer information. A neural net-389 work is then designed and trained by minimizing a specific loss function, that expresses 390 the log-likelihood of the parameters  $\Theta$  describing the probability distribution of desired 391 features  $\mathbf{n}$  that may be either continuous or discrete. 392

# <sup>393</sup> 3 Application to airborne EM data from Morrill, Nebraska

The methodology described above is applied to the inversion of airborne electromagnetic (AEM) data. This inverse problem has been widely studied by deterministic linearized least-squares methods using both a 1D and 3D forward model (Christensen, 2002; Auken & Christiansen, 2004; Viezzoli et al., 2008; Cox et al., 2010; Grayver et al., 2013; Auken et al., 2014).

The full non-linear 1D inverse problem has also been addressed using Markov chain 399 Monte Carlo (McMC) sampling methods, based on for example the reversible-jump sam-400 pling method relying on a prior model representing a 1D layered subsurface (B. J. Mins-401 ley, 2011; B. J. Minsley, Foks, & Bedrosian, 2021; Brodie & Sambridge, 2012). Hansen 402 and Minsley (2019) proposed the use of the extended Metropolis algorithm, also an McMC 403 method, that allows the use of any prior model that can be sampled. The 1D nonlinear 404 inverse EM problem leads to a non-trivial sampling problem, due to the existence of model 405 equivalences (significantly different models lead to the same forward response). Sufficient 406 sampling of the 1D posterior distribution of resistivity values, to obtain a limited set of 407 independent realizations, may require hundreds of thousands of McMC iterations, and 408 hence forward model evaluations. For a single sounding this may take at least 10 min-409 utes per sounding, requiring access to supercomputers for application of real-world data 410 sets (Foks & Minsley, 2020). Hansen (2021) proposed 1D probabilistic inversion based 411 on the extended rejection sampler (using lookup tables, similar to  $[\mathbf{N}^*, \mathbf{M}^*, \mathbf{D}^*]$ ) that rely 412 on the construction of a large sample for the prior along with the forward responses (gen-413 erated once). This is then used to generate independent realizations of the posterior dis-414 tribution numerically more efficiently than is possible using Markov Chain based algo-415 rithms, and at the same time avoids issues related to model equivalences. This sampling 416 approach is used for comparison below. 417

The size of airborne EM surveys is becoming larger, so the use of any of the inver-418 sion methods discussed above will lead to considerable computational demands. Cur-419 rently, two major airborne EM surveys are being carried out. The AusAEM20 project, 420 by Geoscience Australia, is expected to collect around 65,000 flight-line-kilometers of data, 421 422 leading to many hundreds of thousands of EM measurements (Howard, 2020). USGS has collected more than 43,000 flight-line-kilometer data in the Mississippi Alluvial Plain, 423 and another 25,000 flight-line-kilometer is planned for 2021, leading to significantly more 424 than 1,000,000 data points to be inverted in the Mississippi Alluvial Plain (B. J. Mins-425 ley, Rigby, et al., 2021). 426

As an example, we consider the inversion of airborne electromagnetic (AEM) data
from Morrill, Nebraska (Smith et al., 2010; Abraham et al., 2012). We use data at 451
locations, at every 50m along a 22.5 km West-East profile, as also considered in B. J. Minsley (2011). Each observed data set consists of 13 measurements (in-phase and quadrature measurements from 6 pairs of transmitter and receiver coils, as well the measurement altitude).

Three different types of prior models will be defined, that represent different information about the subsurface resistivities  $(\mathbf{m})$  and related (both discrete and contin<sup>435</sup> uous) properties **n** at Morrill. For each of the three prior models considered, a unique <sup>436</sup> posterior probability distribution exists. Various properties of the posterior distribution <sup>437</sup> will be computed using the proposed machine learning method and compared to results <sup>438</sup> obtained from a finite sample of the posterior distributions obtained using the extended <sup>439</sup> rejection sampler with a lookup table of size  $N_T = 2 \cdot 10^6$ .

#### 3.1 A priori models and noise

#### 441 3.1.1 Parameterization

In this example, the subsurface is parameterized into 125 layers of dz = 1 m thickness. Prior models based on up to four sets of parameters,  $\rho(\mathbf{m}, \mathbf{n}_1, \mathbf{n}_2, \mathbf{n}_3)$  are considered.

Resistivity.  $\mathbf{m} = [m_1, m_2, ..., m_{N_M}]$  represents the resistivity of each of the 125 layers.

Layer interface.  $\mathbf{n_1}$  represents the existence of a sharp boundary between two neighboring layers ( $n_{1i} = 0$  when there is no boundary and  $n_{1i} = 1$  in case of a boundary). A sharp boundary is defined when two neighboring resistivity values differ more than 20%.  $\mathbf{n_1}$  refers to 124 discrete parameters and can be directly computed from  $\mathbf{m}$ .

Thickness of highly resistive layer.  $n_2$  represents the cumulative thickness of resistivity values above 225 ohmm. It can be directly computed from m using

$$\mathbf{n}_2 = \sum_{i}^{N_M} dz * I(m_i)$$

where  $I(m_i) = 1$  when  $m_i > 225$  ohmm, and  $I(m_i) = 0$  when  $m_i \le 225$  ohmm.  $\mathbf{n}_2$ refers to a single continuous parameter.

<sup>453</sup> Lithology.  $\mathbf{n_3}$  represents a category ('1', '2', and '3', representing three distinct <sup>454</sup> lithologies) in each layer.  $\mathbf{n_3}$  cannot be computed from  $\mathbf{m}$ , but  $\mathbf{n_3}$  and  $\mathbf{m}$  are linked through <sup>455</sup> a conditional prior distribution  $\rho(\mathbf{m}|\mathbf{n_3})$  (see example below).  $\mathbf{n_3}$  refers to 125 discrete <sup>456</sup> parameters with 3 possible outcomes.

For brevity, all model parameters combined will be referred to as  $\mathbf{p} = [\mathbf{m}, \mathbf{n}_1, \mathbf{n}_2, \mathbf{n}_3]$ . To illustrate the potential of the method 3 different non-Gaussian prior models are considered that vary in complexity and information content.

460

440

#### 3.1.2 Prior information

 $\begin{array}{ll} \rho_A(\mathbf{p}) = \rho_A(\mathbf{m}, \mathbf{n_1}, \mathbf{n_2}), \ a \ uniform \ prior \ model. \ \rho_A(\mathbf{p}) \ represents a \ choice \ of \\ independence \ between \ model \ parameters, \ \rho_A(m_i, m_j) = \rho_A(m_i)\rho_A(m_i) \ \forall (i, j). \ The \ re- \\ sistivity \ of \ each \ resistivity \ model \ parameter \ is \ assumed \ to \ be \ log-uniform \ distributed \ in \\ the \ range \ \mathcal{U}[2,280] \ ohmm. \ This \ is \ the \ least \ informative \ prior \ model \ considered. \ 11 \ in- \\ dependent \ realizations \ of \ \rho_A(\mathbf{m}, \mathbf{n_1}) \ are \ shown \ in \ Figure \ 1a. \end{array}$ 

 $\rho_B(\mathbf{p}) = \rho_B(\mathbf{m}, \mathbf{n_1}, \mathbf{n_2}), \text{ Discrete layered model. } \rho_B(\mathbf{p}) \text{ represents a layered sub-}$ surface consisting of 1 to 8 layers (uniformly distributed), each with a constant resistivity. The resistivity in a specific layer is assumed to be log-uniform distributed in the range  $\mathcal{U}_{\downarrow}[2, 2800] \text{ ohmm.}$ 

A realization  $\mathbf{p}^*$  of  $\rho_B(\mathbf{p})$  is generated by first choosing the number of layers as a random number, Nl, between 1 and 8. Then Nl - 1 layer interfaces are randomly selected from a uniform distribution of  $\mathcal{U}[0, 125]$  m. Then the resistivity within each layer is realized from a uniform distribution  $\mathcal{U}_{\pm}[2, 280]$  ohmm. This type of prior model is similar to the transdimensional prior considered by (B. J. Minsley, 2011). 11 independent realizations of  $\rho_B(\mathbf{m}, \mathbf{n_1})$  are shown in Figure 1c.

476  $\rho_C(\mathbf{p})$ , Trimodal mixture Gaussian.  $\rho_C(\mathbf{p})$  represents a subsurface with three pos-477 sible lithologies ('1', '2' and '3') each with a distinct resistivity distribution. See discus-478 sion about the prior geological knowledge in Morrill in Abraham et al. (2012) and Hansen 479 and Minsley (2019).

To sample  $\rho_C(\mathbf{p}) = \rho_A(\mathbf{m}, \mathbf{n_1}, \mathbf{n_2}, \mathbf{n_3})$ , first a realization of  $\rho_C(\mathbf{n_3})$  is generated 480 as  $\rho_C(\mathbf{n}_3^*)$ , which represents an example of the distribution of the lithologies. This is 481 achieved by generating a realization of a multivariate normal distribution with a Gaus-482 sian type covariance model with a range of 30 m, followed by a simple truncation to ob-483 tain 40% of lithology A, 40% of lithology B, and 20% of lithology C. Then a realization 181 of the resistivity  $\mathbf{m}^*$  is generated, conditional to the lithology type from  $\rho_C(\mathbf{m}|\mathbf{n}_3^*)$ . The resistivity, within each lithology, is generated as a realization of a multivariate normal 486 distribution in  $\log_{10}$ -resistivity space with a range of 30 m, a specific mean,  $m_0$  and stan-487 dard deviation,  $m_{std}$ . For lithology '1',  $m_0 = 1.1$  and  $m_{std} = 0.14$ . For lithology '2', 488  $m_0 = 2$  and  $m_{std} = 0.2$ . For lithology '3',  $m_0 = 2.75$  and  $m_{std} = 0.25$ . Finally,  $n_1^*$ 489 and  $\mathbf{n_2}^*$  are computed from  $\mathbf{m}^*$ . In this way a realization  $\mathbf{p}^* = [\mathbf{m}^*, \mathbf{n_1}^*, \mathbf{n_2}^*, \mathbf{n_3}^*]$  of 490  $\rho_C(\mathbf{p})$  is generated. 11 independent realizations of  $\rho_C(\mathbf{p})$  are shown in Figure 1e. 491

<sup>492</sup>  $\rho_C(\mathbf{p})$  is designed to reflect available information related to the subsurface at Mor-<sup>493</sup> rill (Abraham et al., 2012; Hansen & Minsley, 2019).  $\rho_A(\mathbf{p})$  and  $\rho_B(\mathbf{p})$  are considered <sup>494</sup> here to investigate how the proposed methodology reacts to a uniform (maximum en-<sup>495</sup> tropy) prior such as  $\rho_A(\mathbf{p})$ , and a simple (low entropy) prior as  $\rho_B(\mathbf{p})$ .

#### 3.1.3 Noise

The noise of the EM data is assumed to be independent uncorrelated zero-mean Gaussian noise, with a standard deviation of 5 ppm (parts per million) plus 5 percent noise relative to the noise-free data value, which means the noise depends implicitly on the model. This is the same noise model as considered in previous works on the EM data from Morrill (B. J. Minsley, 2011; Hansen & Minsley, 2019; Hansen, 2021).

#### 502

510

496

## 3.2 Sampling of the posterior distribution

For reference, the extended rejection sampler, with a lookup table of size  $N_T = 5 \cdot 10^6$ , is used to sample the posterior distribution, as detailed in Hansen (2021). 11 independent realizations of the three posterior distributions ( $\sigma_A(\mathbf{p}), \sigma_B(\mathbf{p})$ , and  $\sigma_C(\mathbf{p})$ ) are shown in Figures 1b,d,f.

The goal of the proposed machine learning approach is to directly compute statistical properties of the posterior distribution similar to the same statistical properties obtained from a sample of the posterior using sampling, such as shown in Figures 1b,d,f.

## 3.3 Neural network design

Two fully connected neural networks are designed to allow the characterization of the 1D marginal posterior distribution of continuous and discrete parameters. The input layer, in both cases, consists of the observed data  $\mathbf{d}_{obs}$ , or simulated data with noise. For this specific case, it consists of 13 neurons. 12 neurons refer to the 12 data measurements, and 1 neuron to the altitude measured during recording of data.

The inner network is designed using either 4 or 8 hidden layers depending on the application, each with 40 neurons using the rectified linear activation function (RELU) (Bishop et al., 1995). This inner part of the network needs to be complex enough that the desired mapping can be represented, but simple enough to avoid overfitting, as dis-

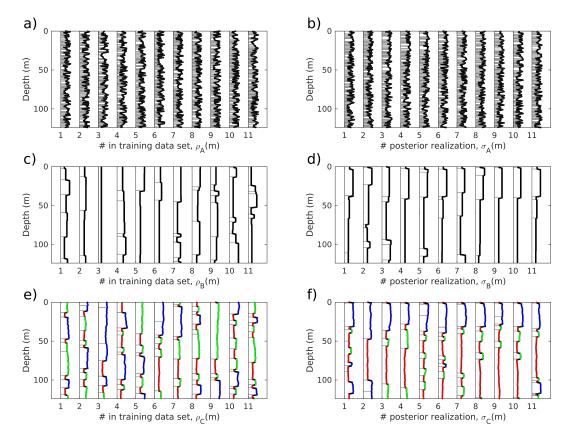


Figure 1. First 11 models from the training data, **T**, for three prior models a)  $\rho_a(\mathbf{m}, \mathbf{n_1})$ , c)  $\rho_b(\mathbf{m}, \mathbf{n_1})$ , and e)  $\rho_c(\mathbf{m}, \mathbf{n_1}, \mathbf{n_2})$ , as well as 11 independent realizations from the posterior distribution obtained for the data at x=15km for b)  $\sigma_b(\mathbf{m}, \mathbf{n_1})$ , d)  $\sigma_d(\mathbf{m}, \mathbf{n_1})$ , and f)  $\sigma_f(\mathbf{m}, \mathbf{n_1}, \mathbf{n_2})$ . Thin horizontal black lines indicate the existence of a layer interface  $(\mathbf{n_1})$ . The thick line indicates variation in resistivity (**m**). In e) and f) the colors of the thick line represent lithology A (red), B(blue), and C (green) when defined.

cussed also by (Meier et al., 2007). Network design is highly problem-dependent, and
 for the present problem, we found this network design provides results on par with, and
 in some cases better than, sampling-based approaches, while at the same time being rel atively easy to optimize.

As discussed, the choice of loss function, and to some extent the activation function, are set by the specific property of the posterior distribution that will be estimated. This leads to two specific types of output layers for regression and classification type problems<sup>1</sup>.

## 3.3.1 Regression type neural network

528

534

543

558

The first neural network type is designed to estimate parameters  $\theta$  of a probability distribution describing the 1D marginal posterior distribution of continuous parameters (such as **m** and **n**<sub>2</sub>). If  $N_{\theta}$  is the number of parameters needed to describe a specific 1D distribution, then in total  $N_{out} = N_{\theta}N_m$  neurons are needed in the output layer if the target is properties of  $\sigma(\mathbf{m})$ , and  $N_{out} = N_{\theta}$  if the target is  $\sigma(\mathbf{n}_2)$ .

## 3.3.2 Classification type neural network

The second neural network type is designed to estimate the posterior probability of possible classes for the discrete type model parameters  $\mathbf{n_1}$  and  $\mathbf{n_3}$ , i.e. of  $\sigma(\mathbf{n_1}) \sigma(\mathbf{n_3})$ .

If the goal is to estimate the 1D marginal distribution of a discrete parameter with  $N_{cat}$  possible outcomes, this can be achieved by selecting an output layer with  $N_{out} =$   $N_m$  when  $N_{cat} = 2$  (using a sigmoid activation function), and  $N_{out} = N_{cat}N_m$  when  $N_{cat} > 2$  (using the softmax activation function). As discussed above, using the crossentropy loss function, Eqn. 18, will lead to direct estimation of the 1D posterior marginal probabilities in this case.

#### 3.4 Network training

Using the prior models, the nonlinear forward model, and the noise model, a train-544 ing data set of size  $N_T = 5 \cdot 10^6$  is constructed (one for each type of prior model) and 545 used for training. Both networks are trained using 67% of the training data set, while 546 33% is reserved for validation. In both cases, the loss function is minimized using the 547 Adam optimizer (Kingma & Ba, 2014) using a learning rate of 0.001, for a maximum of 548 2000 epochs. Early stopping is utilized which stops the training if the loss function eval-549 uated on the validation data does not decrease for 50 epochs. This is done to avoid over-550 fitting, where the loss on the training data will decrease, but where the loss on the val-551 idation data increases. TensorFlow with Keras and TensorFlow-probability have been 552 used to implement and train the neural networks (Abadi et al., 2015; Chollet, 2015; Dil-553 lon et al., 2017). 554

The two considered networks, and the training of the networks, only differ concerning the definition of the output layer (the number of nodes and activation function), the choice of loss function, and the chosen number of hidden layers.

# 3.5 Estimation of properties of $\sigma(m)$

First, properties related to the posterior distribution of resistivity,  $\sigma(\mathbf{m})$ , are considered.

<sup>&</sup>lt;sup>1</sup> Example implementations of these two types of neural networks can be found at http://github.com/ cultpenguin/ip\_and\_ml/.

# 3.5.1 Estimation of mean and standard deviation of $\sigma(\mathbf{m})$

A neural network is set up and trained to estimate the pointwise mean and standard deviation of  $\sigma(\mathbf{m})$ , using 8 hidden layers, by minimizing the loss function in Eqn 14.

Figures 2a-e shows the pointwise mean of the posterior distribution  $\sigma_C(\mathbf{m})$  obtained using the machine learning approach with a training data set of size  $N_T = [10^3, 10^4, 10^5, 10^6, 5 \cdot 10^6]$ , compared to the same statistics computed from a sample of the posterior obtained using the sampling method, Figure 2f. The corresponding standard deviation is shown in Figures 2g-l.

It is clear from Figures 2a,g that using  $N_T = 10^3$  provides very poor results, as compared to the results obtained using sampling, Figures 2e,j. But even using  $N_T =$  $10^4$  leads to results close to the sampling-based results. The changes in predicted mean and standard deviation become smaller as  $N_T$  increases, with only very subtle changes between the use of  $N_T = 10^6$  and  $N_T = 5 \cdot 10^6$ .

One notable difference when comparing Figures 2e,k ( $N_T = 5 \cdot 10^6$ ) and 2f,l (sampling), is that sampling results in more small scale variability in the estimated parameters, as opposed to the more smooth result obtained using machine learning. The reason is simply that the sampling-based approach is based on inferring the statistics from a finite-sized sample of the posterior, whereas in the machine learning approach these statistics are estimated directly.

Figure 3 shows a comparison between the posterior mean and standard deviation obtained using the sampling approach and using the machine learning approach ( $N_T = 5 \cdot 10^6$ ), for  $\sigma_A(\mathbf{m})$  (Figures 3a-d) and  $\sigma_B(\mathbf{m})$  (Figures 3e-h) respectively.

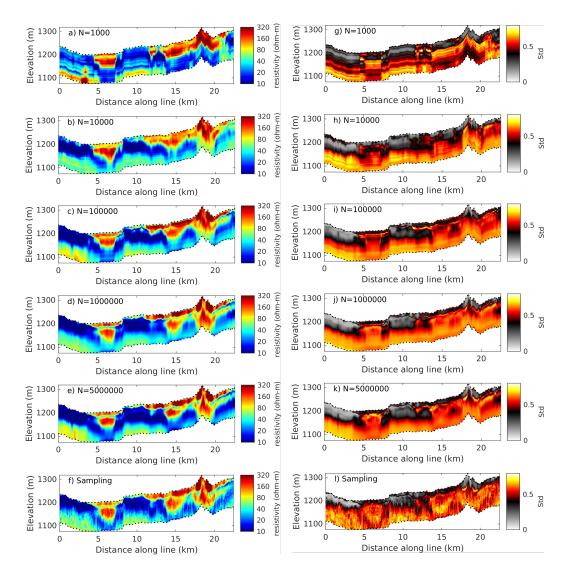
 $\rho_A(\mathbf{m})$  refers to the least informed prior model, and hence one should expect the least resolution in the corresponding posterior distribution. This is what can be seen in results from both the machine learning and the sampling approach, Figures 3a-d, where only the resistivities at the top of the model are resolved.

<sup>588</sup> While  $\rho_B(\mathbf{m})$  is somewhat simpler than  $\rho_C(\mathbf{m})$ , the mean of the corresponding pos-<sup>589</sup> terior distribution are rather similar, Figures 2e and 3f, with the largest difference re-<sup>590</sup> lated to the posterior standard deviation, Figures 2k and 3h.

A key point from Figures 2 and 3 is that the use of the machine learning based approach seems to provide results at least on par with the results obtained using sampling, when the goal is to estimate the mean and standard deviation of the (non-Gaussian) posterior distribution. This is the case using both informed and uninformed prior models.

3.5.1.1 Computational efficiency Figure 4 shows the training and validation loss, 595 and computation time<sup>2</sup> needed to train the neural networks for the results presented in 596 Figure 2. The training time increases with the size of the training data set,  $N_T$ . Both 597 training and validation loss are reduced when  $N_T$  increases. It is also clear that the rel-598 ative difference in loss decreases when comparing the use of  $N_T = 10^5$  to  $N_T = 5$ . 599 10<sup>6</sup>, to when comparing the use of  $N_T = 10^3$  to  $N_T = 10^5$ . Hence, using  $N_T > 10^5$ 600 leads to a substantial longer training time, but only to a minor loss reduction. In addi-601 tion, and as expected, when  $N_T$  increases the validation loss seems to match the train-602 ing loss increasingly well, which indicates that there is no problem with overfitting. 603

 $<sup>^2\,\</sup>mathrm{a}$  workstation with an Intel Core(TM) i7-8700K CPU, Nvidia RTX 3090 GPU, and 64 Gb RAM was used



**Figure 2.** Pointwise mean (a-f) and standard deviation (g-l) obtained from  $\sigma_C(\mathbf{m})$  obtained using machine learning based on a training data set of size  $10^3$  (a,g),  $10^4$  (b,h),  $10^5$  (c,i),  $10^6$  (d,j),  $5 \cdot 10^6$  (e,k), and using the extended rejection sampler (f,l).

Once set up and trained, the prediction of the network is very fast. For all the networks presented above, the prediction time for all 451 data locations was less than 5ms. This means that more than 100,000 soundings can be analyzed per second.

607

# 3.5.2 Estimation of multiple 1D properties of $\sigma(m_i)$

As described above, any parameter of a probability distribution for which a loss 608 function can be described through Eqn. 11 can be estimated using the machine learn-609 ing method. To demonstrate this, 4 independent networks have been trained to estimate 610 properties ( $\Theta$ ) of the 1D marginal posterior distribution  $\sigma(m_i)$  given by a) a normal dis-611 tribution (Eqn. 12, as in Figure 2), b) a generalized normal distribution (Eqn. 15), c) 612 a mixture distribution based on two Gaussian distributions (Eqn. 15), and d) a mixture 613 distribution based on three Gaussian distributions (Eqn. 15). The loss functions used 614 are the negative log-likelihood of the probability distribution in Eqns. 12, 15, and 16 re-615 spectively. 616

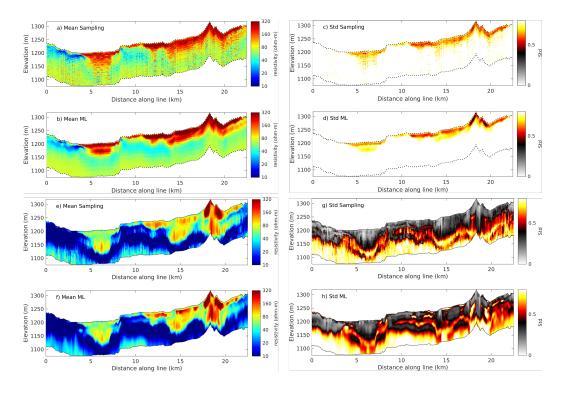


Figure 3. Pointwise mean (a,b) and standard deviation (c,d) obtained from  $\sigma_A(\mathbf{m})$  obtained using the extended rejection sampler (a,c) and machine learning (b,d) based on a training data set of size  $N_T = 5 \cdot 10^6$ . e-h) As a-d) but for  $\sigma_B(\mathbf{m})$ .

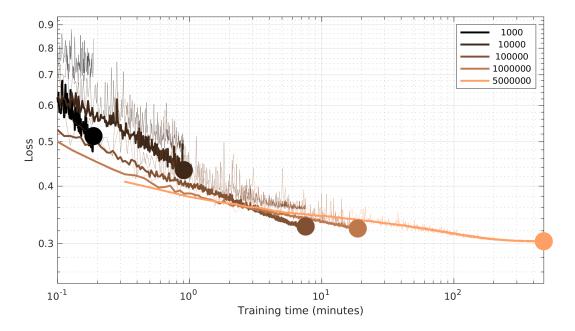


Figure 4. Training (thick lines) and validation (thin lines) loss as a function of training time for  $N_t = [10^3, 10^4, 10^5, 10^6, 5 \cdot 10^6]$ .

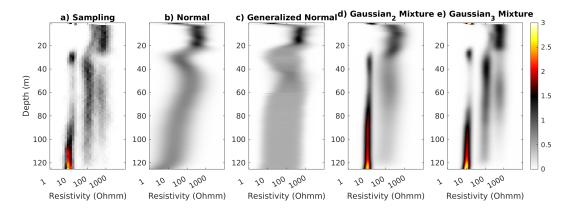


Figure 5. 1D posterior probability density with depth using data at X=6.2 km a) obtained using sampling followed by computation of the marginal posterior probability, and constructed from statistical properties inferred for a b) normal distribution, c) generalized normal distribution, d)-e) a mixture model based on 2 and 3 1D normal distributions. Obtained using  $N_T = [5 \cdot 10^6]$ .

The number of parameters to estimate for the 4 cases, and hence neurons in the output layer, are  $N_{\theta} = [2 * N_m, 3 * N_m, 2 * N_m * N_c, 3 * N_m * N_c] = [250, 375, 750, 1125],$ where  $N_c$  is the number of distributions in the mixture model.

Figure 5a shows the posterior 1D marginal distribution of resistivity values obtained using sampling, based on a finite set of realizations, obtained at x=15 km. One can clearly identify a bimodal to trimodal distribution at depth representing the three possible lithologies from the prior model  $\rho_C(\mathbf{m})$  with different resistivity values.

Figures 5b-e, shows the probability distributions representing the estimated statistical properties of the 4 considered distributions. These distributions do not represent assumptions about the posterior distribution (which can be arbitrarily complex) but reflect example statistical properties that one might calculate from a sample of the posterior.

If the goal is to compute a representation of the 1D posterior marginal distribution, as considered by (Meier et al., 2007; Shahraeeni & Curtis, 2011), then care should be taken to use a parameterization for the chosen 1D distribution complex enough to describe the variability of the posterior. From Figure 5 it is evident that only in case using the mixture model with 3 Gaussian distributions, the estimated marginal probability density represents the actual 1D marginal posterior distribution well.

The statistical properties of the posterior distribution which it is relevant to compute for a specific inverse problem, are naturally problem-dependent. This example nonetheless demonstrates that the machine learning methodology is capable of estimating parameters of different types of probability distributions, for which a probability density, and hence the corresponding loss function, can be computed.

# $_{640}$ 3.6 Estimation of properties of $\sigma(n_1)$

 $\sigma(\mathbf{n_1})$  refers to the existence (or lack of) a layer interface, which can be formulated as a binary classification problem. Therefore, a classification type network is constructed using a sigmoid activation function, and Eqn. 18 as the loss function. 4 hidden layers are used. Figures 6a and 6c refer to the pointwise posterior probability of locating a layer interface, as computed from a sample from the posterior distribution of  $\sigma_B(\mathbf{n_1})$  and  $\sigma_C(\mathbf{n_1})$ . The corresponding results obtained as the output of a trained neural network based on a training data set of size  $N_T = 5 \cdot 10^6$  are shown in Figures 6b and 6d. The prior probability of a layer interface is around 0.1, and hence a posterior probability of 0.25 is indicative of a layer interface.

The results using sampling and the machine learning approach are in both cases very similar with a bit more variability in the results obtained using sampling, due to the use of a finite-sized sample of the posterior distribution.

#### 654

# 3.7 Estimation of properties of $\sigma(n_2)$

We consider the simpler problem of inferring information about a single continuous parameter,  $n_2$ , representing the cumulative thickness of layers with a resistivity above 225 ohmm. The same neural network as considered above to estimate properties related to **m** is used here, except that here only 4 hidden layers are used.

Figure 7 shows the mean of  $\sigma_C(\mathbf{n_2})$  (black line), as well as the probability distribution reflecting the mean and standard deviation estimated using the machine learning approach for  $N_T = [10^3, 10^4, 10^5, 10^6, 5 \cdot 10^6]$  in figures 7a-e. The mean computed using the machine learning approach compares well to the mean obtained using sampling methods for  $N_T \ge 10^5$ .

# $_{664}$ 3.8 Estimation of properties of $\sigma(n_3)$

Finally, we consider the discrete parameter  $\mathbf{n_3}$  which refer to lithology type, which can be of type '1', '2' and '3'. The outcome for each model parameter is then a multiclass (three classes) classification problem. Therefore, a classification type network is constructed using a softmax activation function, and the loss function in Eqn. 18. 4 hidden layers are used.

Figures 8a,c,e show the posterior probability for each of the three classes obtained using sampling, while Figures 8b,d,f show the corresponding results obtained by evaluating the trained network. Except for some small-scale variations in the sampling results, due to using a finite sample size, the obtained posterior statistics are strikingly similar.

# 675 4 Discussion

A typical application of probabilistic inversion is to use some sampling method to generate a large sample from the posterior distribution. Then some appropriate statistic, computed from the sample of the posterior distribution, is chosen and visualized.

The theory presented above proposes how one can construct a neural network that can directly estimate any statistical property of the posterior distribution (for discrete or continuous parameters) for which a probability distribution can be evaluated, without ever generating realizations of the posterior distribution. This can be achieved by the following steps:

- 1. Construct a training data set, in the style of Devilee et al. (1999),  $\mathbf{T}^* = [\mathbf{N}^*, \mathbf{D}^*_{sim}]$ , where  $\mathbf{N}^*$  represents a set of features/properties of interest, and  $\mathbf{D}^*_{sim}$  represents a corresponding set of simulated data with noise, using both the forward and the noise model.
- <sup>688</sup> 2. Design a neural network whose output layer represents the relevant statistical pa-<sup>689</sup> rameters  $\Theta$  of the posterior distribution  $\sigma(\mathbf{n})$  of interest.

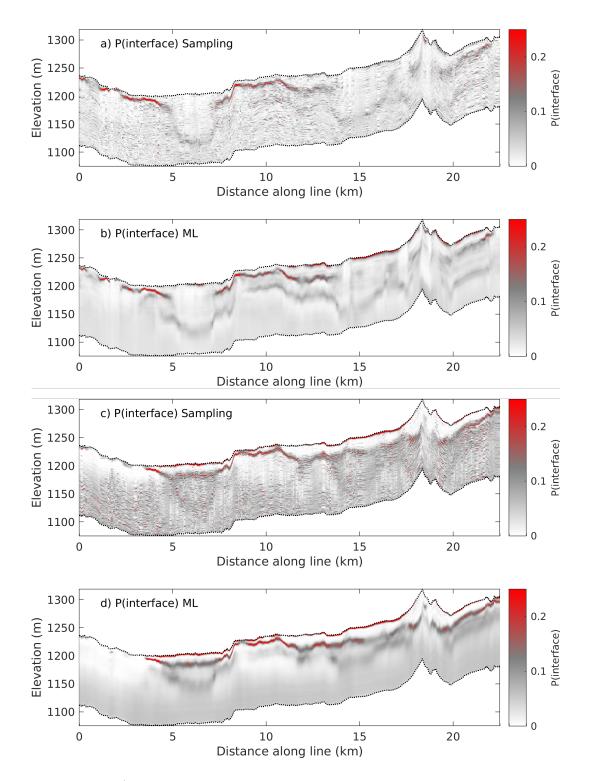


Figure 6. a-b) Posterior probability of a layer interface obtained using extended rejection sampling (a), and machine learning (b), for  $\sigma_B(\mathbf{n}_1)$ . c-d) Posterior probability of a layer interface obtained using extended rejection sampling (c), and machine learning (d), for  $\sigma_C(\mathbf{n}_1)$ . Obtained using  $N_T = [5 \cdot 10^6]$ .

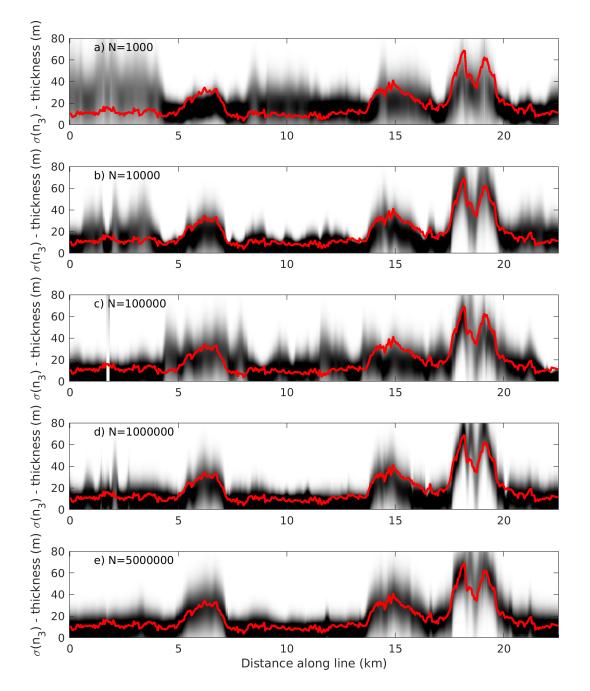


Figure 7. Mean of the posterior distributions  $\sigma_C(\mathbf{n_3})$  estimated using sampling (red line) compared with the estimated 1d normal mean and standard deviation of  $\sigma_C(\mathbf{n_3})$  plotted as probability density in grayscale, estimated using the machine learning approach using a training data set of size  $N_T = [10^3, 10^4, 10^5, 10^6, 5 \cdot 10^6]$  in a)-e).

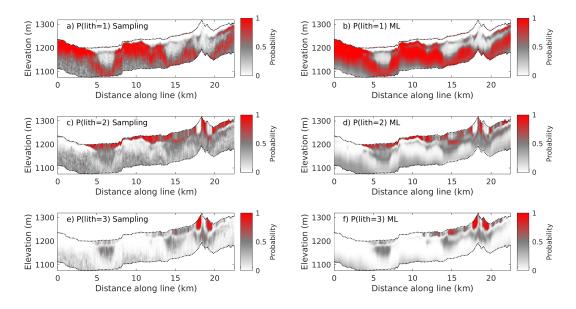


Figure 8. Posterior probability of lithology, using a-c) sampling and d-f) machine learning for  $\sigma_C(\mathbf{n}_3)$ .

3. Train the neural network by minimizing a loss function that is the negative loglikelihood of the probability density,  $f(\Theta)$ , whose properties one wishes to estimate.

Practical application of the methodology requires a) a neural network structure complex enough to be able to estimate the mapping  $\mathbf{d}_{sim}^* \mapsto \Theta$ , and b) a training data set large enough to allow the mapping to be inferred.

The methodology was applied and demonstrated in a case study using airborne EM 696 data from Morrill, Nebraska. Several (uninformed to more informed) prior models were 697 considered, describing both subsurface resistivity (a continuous parameter,  $\mathbf{m}$ ) and lithol-698 ogy (a discrete parameter,  $n_3$ ) and the considered forward problem was nonlinear. In 699 addition, the method was used to estimate posterior statistics of low-dimensional fea-700 tures of the prior models, such as the existence of a layer interface,  $\mathbf{n_1}$ , and the thick-701 ness of layers with resistivity above 225 ohmm,  $n_2$ . Results showed that using a train-702 ing data set of size  $N_T > 10^5$  in this case leads to a trained neural network that pro-703 vides estimates of posterior statistics similar to those obtained using sampling methods, 704 using a fraction of the computational power (about 5ms per sounding). 705

# 4.1 Limitations

706

The proposed method does not generate realizations of the posterior distribution,
as do other sampling-based methods (B. J. Minsley, 2011; Brodie & Sambridge, 2012;
Hansen & Minsley, 2019; Hansen, 2021). Instead, statistics of the posterior distribution
for features of interest are estimated directly by applying a trained neural network.

In some use cases, one may actually need the realizations, for example to propagate flow responses from of a set of realizations from the posterior representing hydraulic parameters, (Vilhelmsen et al., 2019). But, in many applications, where one is primarily interested in some statistical parameter describing the posterior, such as the posterior probability of a lithology type, the presented methodology may be useful. The key practical difference to using sampling methods is that one has to identify the feature one is interested in and specify an appropriate loss function before running the inversion. Whereas using sampling methods to sample the posterior, one can convert the realizations of the posterior into a specific feature, and perform the posterior analysis, after the sampling algorithm has run.

In the example application we adopted a widely used uncorrelated Gaussian noise model. In practice, real data are often affected by correlated noise (Hansen et al., 2014; Hauser et al., 2015; Bai et al., 2021). While in principle any noise model can be handled by the proposed methodology, as long as realizations of the noise can be generated, it remains to be tested how well the methodology works with more complex noise models.

The methodology is particularly promising for localized inverse problems, where 726 the trained neural network can be set up and trained once, but applied many times. It 727 is less obviously suited to 3D inversions with very large model dimensions because 1) con-728 struction of an adequately large training data set will be difficult and CPU intensive, 2) 729 solving the 3D forward problem may be CPU intensive, and 3) it may be very difficult 730 to train a neural network with millions of parameters in the output layer. While the method 731 appears to work well for the AEM case considered, it remains to be seen how the method 732 performs for other, possibly more nonlinear, inverse problems. 733

# 4.2 Potential

734

The immediate appeal of the proposed methodology is that it leads to fast predic-735 tion times. One can get similar results, but much faster, compared with using sampling-736 based methods to analyze the posterior distribution. The presented method is faster than 737 linearized least squares based deterministic inversion of EM data (which use less than 738 a second CPU time per 4 soundings), which have been widely used for inversion of large 739 surveys (Auken et al., 2017; B. Minsley et al., 2021) because they require much less com-740 putational resources than sampling-based methods. With the computational efficiency 741 of the proposed method, the computational benefits of linearized methods are no longer 742 so substantial that one should ignore the benefits of using the probabilistic methods that 743 allow the use of site-specific prior information, a non-linear forward model, and full ex-744 ploration of the space of uncertainty. 745

The more general appeal is that the proposed methodology allows the use of in prin-746 ciple arbitrarily complex prior models. The only requirement is that one must be able 747 to generate independent realizations of the prior model. This allows an end-user to ac-748 tively choose a prior model based on available information, as opposed to being forced 749 to use the prior assumptions implicit in most available inversion algorithms, such as the 750 assumptions of a layered subsurface (B. J. Minsley, Foks, & Bedrosian, 2021) or a Gaus-751 sian type smooth prior (Auken & Christiansen, 2004). The prior can be constructed ac-752 cording to site-specific information, and posterior statistics can be estimated of any pa-753 rameter that can be computed from the prior model, as illustrated by the parameters 754  $\mathbf{n}_1$  and  $\mathbf{n}_2$  in the case study 755

The main challenge then becomes the construction of realistic prior models that represent geological realistic information as well as realistic noise models.

# 758 5 Conclusions

A simple, yet powerful, approach to probabilistic inversion has been proposed. Its application requires that one can simulate sets of examples capturing the known information. That is 1) sample from an arbitrarily complex prior model, 2) solving the forward problem, and 3) adding realistic noise to the simulated data. From each of these sets of models and data, a set of corresponding features related to the model parameters can be obtained. Together these represent, up to the limit of the finite set of models, all known information about these features of interest.

From such sets of features and corresponding noisy input data, a neural network can be used to estimate statistical properties of the posterior distribution directly, by training the network to minimize an appropriate loss function. This provides the ability to carry out a fast and accurate estimation of relevant posterior statistics given an observed dataset.

A case study of the methodology applied to a nonlinear probabilistic inversion of 771 EM data demonstrates it is possible to directly obtain posterior statistics similar to those 772 obtained using sampling methods, using a fraction of the computation time. This ap-773 proach allows the use and testing of multiple prior models, and to consider multiple fea-774 tures related to the prior distributions, in a fully probabilistic setting using only mod-775 est computational resources. The method has most appeal for localized inverse problems, 776 where the same trained neural network can be applied to many datasets with little com-777 putational effort. 778

# 779 Data Availability Statement

The airborne EM data used in this study is freely available (Smith et al., 2010) and can be accessed through https://doi.org/10.3133/ofr20101259. Training data sets and python notebooks for training and prediction will be made available upon publication at https://github.com/cultpenguin/probabilistic-inverse-problems\_and\_ml

# 784 Acknowledgments

This work is funded by the Danish Free Research Council, project 717-00160B.

# 786 **References**

- Abadi, M., Agarwal, A., Barham, P., Brevdo, E., Chen, Z., Citro, C., ... Zheng, X.
   (2015). TensorFlow: Large-scale machine learning on heterogeneous systems.
   Retrieved from https://www.tensorflow.org/ (Software available from
- <sup>790</sup> tensorflow.org)
- Abraham, J. D., Cannia, J. C., Bedrosian, P. A., Johnson, M. R., Ball, L. B., &
   Sibray, S. S. (2012). Airborne electromagnetic mapping of the base of aquifer
   *in areas of western nebraska*. U.S. Geological Survey Scientific Investigations
   Report 2011-5219, 38 p.
- Ardizzone, L., Kruse, J., Wirkert, S., Rahner, D., Pellegrini, E. W., Klessen, R. S.,
   ... Köthe, U. (2018). Analyzing inverse problems with invertible neural networks. arXiv preprint arXiv:1808.04730.
- Auken, E., Boesen, T., & Christiansen, A. V. (2017). A review of airborne electromagnetic methods with focus on geotechnical and hydrological applications from 2007 to 2017. Advances in Geophysics, 58, 47–93.
- Auken, E., & Christiansen, A. V. (2004). Layered and laterally constrained 2d inversion of resistivity data. *Geophysics*, 69(3), 752–761.
- Auken, E., Christiansen, A. V., Kirkegaard, C., Fiandaca, G., Schamper, C.,
  Behroozmand, A. A., ... others (2014). An overview of a highly versatile
  forward and stable inverse algorithm for airborne, ground-based and borehole
  electromagnetic and electric data. *Exploration Geophysics*, 46(3), 223–235.
- Bai, P., Vignoli, G., Andrea, V., Jouni, N., & Vacca, G. (2020). (quasi-) real-time
   inversion of airborne time-domain electromagnetic data via artificial neural
   network. *Remote Sensing*.
- Bai, P., Vignoli, G., & Hansen, T. M. (2021). 1d stochastic inversion of airborne
   time-domain electromagnetic data with realistic prior and accounting for the

812	forward modeling error. Remote Sensing, $13(19)$ , $3881$ .
813	Bishop, C. M., et al. (1995). Neural networks for pattern recognition. Oxford univer-
814	sity press.
815	Blei, D. M., Kucukelbir, A., & McAuliffe, J. D. (2017). Variational in erence:
816	A review for statisticians. Journal of the American statistical Association,
817	112(518), 859-877.
818	Bording, T. S., Asif, M. R., Barfod, A. S., Larsen, J. J., Zhang, B., Grombacher,
819	D. J., others (2021). Machine learning based fast forward modelling of
820	ground-based time-domain electromagnetic data. Journal of Applied Geo-
821	physics, 104290.
822	Bosch, M., Mukerji, T., & Gonzalez, E. F. (2010). Seismic inversion for reservoir
823	properties combining statistical rock physics and geostatistics: A review. Geo-
824	physics, 75(5), 75A165–75A176.
825	Brodie, R. C., & Sambridge, M. (2012). Transdimensional Monte Carlo inversion of
826	AEM data. ASEG Extended Abstracts, 2012(1), 1–4.
827	Chollet, F. (2015). Keras. https://github.com/fchollet/keras. GitHub.
828	Christensen, N. B. (2002). A generic 1-D imaging method for transient electromag-
829	netic data. $Geophysics$ , $67(2)$ , $438-447$ .
830	Constable, S. C., Parker, R. L., & Constable, C. G. (1987). Occam's inversion: A
831	practical algorithm for generating smooth models from electromagnetic sound-
832	ing data. $Geophysics, 52(3), 289-300.$
833	Conway, D., Alexander, B., King, M., Heinson, G., & Kee, Y. (2019). Inverting
834	magnetotelluric responses in a three-dimensional earth using fast forward ap-
835	proximations based on artificial neural networks. Computers $\mathcal{E}$ geosciences,
836	<i>127</i> , 44–52.
837	Cox, L. H., Wilson, G. A., & Zhdanov, M. S. (2010). 3D inversion of airborne elec-
838	tromagnetic data using a moving footprint. Exploration Geophysics, 41(4),
839	250-259.
839 840	250–259. Devilee, R., Curtis, A., & Rov-Chowdhury, K. (1999). An efficient, probabilistic
840	Devilee, R., Curtis, A., & Roy-Chowdhury, K. (1999). An efficient, probabilistic
	Devilee, R., Curtis, A., & Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave
840 841	Devilee, R., Curtis, A., & Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. <i>Journal of Geophysical Research: Solid</i>
840 841 842	Devilee, R., Curtis, A., & Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. <i>Journal of Geophysical Research: Solid Earth</i> , 104 (B12), 28841–28857.
840 841 842 843	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of</li> </ul>
840 841 842 843 844	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. <i>Journal of Geophysical Research: Solid Earth</i>, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural net-</li> </ul>
840 841 842 843 844 845	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> </ul>
840 841 842 843 844 845 846	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. <i>Journal of Geophysical Research: Solid Earth</i>, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural net-</li> </ul>
840 841 842 843 844 845 846 847	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D.,</li> </ul>
840 841 842 843 844 845 846 847 848	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> </ul>
840 841 842 843 844 845 846 847 848 849	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195 (1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint</li> </ul>
840 841 842 843 844 845 846 847 848 849 850	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> </ul>
840 841 842 843 844 845 846 847 848 849 850	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network.</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853 854	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853 854	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853 854 855	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853 854 855 856	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution of tomographic inverse problems. Geophysical Journal International, 216(2), 1344–1363.</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853 854 855 855 856	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution of tomographic inverse problems. Geophysical Journal International, 216(2), 1344–1363.</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853 854 855 856 857 858	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution of tomographic inverse problems. Geophysical Journal International, 216(2), 1344–1363.</li> <li>Foks, N., &amp; Minsley, B. (2020). Geophysical Bayesian inference in Python (GeoBiPy). https://github.com/usgs/geobipy. GitHub. doi: 10.5066/P9K3YH9O</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 855 855 855 855 855 855 855 855	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution of tomographic inverse problems. Geophysical Journal International, 216(2), 1344–1363.</li> <li>Foks, N., &amp; Minsley, B. (2020). Geophysical Bayesian inference in Python (GeoBiPy). https://github.com/usgs/geobipy. GitHub. doi: 10.5066/P9K3YH9O</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853 854 855 856 855 856 857 858 859 860	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution of tomographic inverse problems. Geophysical Journal International, 216(2), 1344–1363.</li> <li>Foks, N., &amp; Minsley, B. (2020). Geophysical Bayesian inference in Python (GeoBiPy). https://github.com/usgs/geobipy. GitHub. doi: 10.5066/P9K3YH9O</li> <li>Geman, S., &amp; Geman, D. (1984). Stochastic relaxation, gibbs distributions, and</li> </ul>
840 841 842 843 844 845 846 847 848 849 850 851 852 853 854 855 856 857 858 859 860 861	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution of tomographic inverse problems. Geophysical Journal International, 216(2), 1344–1363.</li> <li>Foks, N., &amp; Minsley, B. (2020). Geophysical Bayesian inference in Python (GeoBiPy). https://github.com/usgs/geobipy. GitHub. doi: 10.5066/P9K3YH9O</li> <li>Geman, S., &amp; Geman, D. (1984). Stochastic relaxation, gibbs distributions, and the bayesian restoration of images. IEEE Transactions on pattern analysis and</li> </ul>
<ul> <li>840</li> <li>841</li> <li>842</li> <li>843</li> <li>844</li> <li>845</li> <li>846</li> <li>847</li> <li>848</li> <li>849</li> <li>850</li> <li>851</li> <li>855</li> <li>856</li> <li>857</li> <li>858</li> <li>859</li> <li>860</li> <li>861</li> <li>862</li> <li>863</li> </ul>	<ul> <li>Devilee, R., Curtis, A., &amp; Roy-Chowdhury, K. (1999). An efficient, probabilistic neural network approach to solving inverse problems: inverting surface wave velocities for eurasian crustal thickness. Journal of Geophysical Research: Solid Earth, 104 (B12), 28841–28857.</li> <li>de Wit, R. W., Valentine, A. P., &amp; Trampert, J. (2013). Bayesian inference of earth's radial seismic structure from body-wave traveltimes using neural networks. Geophysical Journal International, 195(1), 408–422.</li> <li>Dillon, J. V., Langmore, I., Tran, D., Brevdo, E., Vasudevan, S., Moore, D., Saurous, R. A. (2017). Tensorflow distributions. arXiv preprint arXiv:1711.10604.</li> <li>Earp, S., &amp; Curtis, A. (2020). Probabilistic neural network-based 2d travel-time tomography. Neural Computing and Applications, 32(22), 17077–17095.</li> <li>Earp, S., Curtis, A., Zhang, X., &amp; Hansteen, F. (2020, 08). Probabilistic neural network tomography across Grane field (North Sea) from surface wave dispersion data. Geophysical Journal International, 223(3), 1741-1757. Retrieved from https://doi.org/10.1093/gji/ggaa328</li> <li>Fichtner, A., Zunino, A., &amp; Gebraad, L. (2018). Hamiltonian monte carlo solution of tomographic inverse problems. Geophysical Journal International, 216(2), 1344–1363.</li> <li>Foks, N., &amp; Minsley, B. (2020). Geophysical Bayesian inference in Python (GeoBiPy). https://github.com/usgs/geobipy. GitHub. doi: 10.5066/P9K3YH9O</li> <li>Geman, S., &amp; Geman, D. (1984). Stochastic relaxation, gibbs distributions, and the bayesian restoration of images. IEEE Transactions on pattern analysis and machine intelligence(6), 721–741.</li> </ul>

Grayver, A. V., Streich, R., & Ritter, O. (2013).Three-dimensional parallel dis-867 tributed inversion of CSEM data using a direct forward solver. Geophysical 868 Journal International, 193(3), 1432–1446. 869 Green, P. J. (1995). Reversible jump Markov chain Monte Carlo computation and 870 Bayesian model determination. *Biometrika*, 82(4), 711–732. 871 Hansen, T. M. (2021).Efficient probabilistic inversion using the rejection sam-872 pler—exemplified on airborne EM data. Geophysical Journal International, 873 224(1), 543-557.874 Hansen, T. M., Cordua, K. C., & Mosegaard, K. (2012). Inverse problems with non-875 trivial priors - efficient solution through sequential Gibbs sampling. Computa-876 tional Geosciences, 16(3), 593-611. doi: 10.1007/s10596-011-9271-1 877 Hansen, T. M., & Cordua, K. S. (2017). Efficient monte carlo sampling of inverse 878 problems using a neural network-based forward—applied to gpr crosshole 879 traveltime inversion. Geophysical Journal International, 211(3), 1524–1533. 880 Hansen, T. M., Cordua, K. S., Jacobsen, B. H., & Mosegaard, K. (2014). Accounting 881 for imperfect forward modeling in geophysical inverse problems—exemplified 882 for crosshole tomography. Geophysics, 79(3), H1–H21. 883 Hansen, T. M., Cordua, K. S., Looms, M. C., & Mosegaard, K. (2013, 03). SIPPI: 884 a Matlab toolbox for sampling the solution to inverse problems with com-885 plex prior information: Part 1, methodology. Computers & Geosciences, 52, 886 470-480. doi: 10.1016/j.cageo.2012.09.004 887 Hansen, T. M., Cordua, K. S., Zunino, A., & Mosegaard, K. (2016).Probabilit-888 sic Integration of Geo-Information. In M. Moorekamp (Ed.), Joint inversion 889 (Vol. 218). John Wiley & Sons. 890 Hansen, T. M., & Minsley, B. J. (2019). Inversion of airborne EM data with an ex-891 plicit choice of prior model. Geophysical Journal International, 218(2), 1348-892 1366.893 Hansen, T. M., Mosegaard, K., & Cordua, K. C. (2008).Using geostatistics to 894 In J. M. Ortiz & describe complex a priori information for inverse problems. 895 X. Emery (Eds.), Viii international geostatistics congress (Vol. 1, p. 329-338). 896 Mining Engineering Department, University of Chile. 897 Hastings, W. K. (1970). Monte Carlo sampling methods using Markov chains and 898 their applications. *Biometrika*, 57(1), 97. 899 Hauser, J., Gunning, J., & Annetts, D. (2015). Probabilistic inversion of airborne 900 electromagnetic data under spatial constraints. Geophysics, 80(2), E135– 901 E146. 902 Hornik, K., Stinchcombe, M., & White, H. (1990). Universal approximation of an 903 unknown mapping and its derivatives using multilayer feedforward networks. 904 Neural networks, 3(5), 551-560. 905 Howard, D. (2020). Geological survey of western australia: Ausaem20-wa project. 906 Preview, 2020(205), 18-18.907 Khoshkholgh, S., Zunino, A., & Mosegaard, K. (2022).Full-waveform inversion 908 by informed-proposal monte carlo. Geophysical Journal International, 230(3), 909 1824 - 1833.910 Kingma, D. P., & Ba, J. (2014). Adam: A method for stochastic optimization. arXiv 911 preprint arXiv:1412.6980. 912 Köpke, C., Irving, J., & Elsheikh, A. H. (2018).Accounting for model error in 913 Bayesian solutions to hydrogeophysical inverse problems using a local basis 914 approach. Advances in Water Resources, 116, 195–207. 915 Laloy, E., Hérault, R., Jacques, D., & Linde, N. (2018). Training-image based geo-916 statistical inversion using a spatial generative adversarial neural network. Wa-917 ter Resources Research, 54(1), 381-406. 918 Laloy, E., & Vrugt, J. A. (2012). High-dimensional posterior exploration of hydro-919 logic models using multiple-try dream (zs) and high-performance computing. 920 Water Resources Research, 48(1). 921

- Madsen, R. B., & Hansen, T. M. (2018). Estimation and accounting for the mod eling error in probabilistic linearized amplitude variation with offset inversion.
   *Geophysics*, 83(2), N15–N30.
- Malinverno, A. (2002). Parsimonious Bayesian Markov chain Monte Carlo inversion in a nonlinear geophysical problem. *Geophysical Journal International*, 151(3), 675–688.
- Meier, U., Curtis, A., & Trampert, J. (2007). Global crustal thickness from neural network inversion of surface wave data. *Geophysical Journal International*, 169(2), 706–722.
- Menke, W. (2012). Geophysical data analysis: Discrete inverse theory (Vol. 45).
   Academic Press.
- Metropolis, N., Rosenbluth, M., Rosenbluth, A., Teller, A., & Teller, E. (1953).
   Equation of state calculations by fast computing machines. J. Chem. Phys., 21, 1087–1092.

936

937

938

939

940

941

942

943

944

949

950

951

- Minsley, B., Rigby, J., James, S., Burton, B., Knierim, K., Pace, M., ... Kress, W. (2021). Closing a scale-gap in earth observation using regional-scale airborne geophysics in the lower mississippi valley (Tech. Rep.). Copernicus Meetings.
- Minsley, B. J. (2011). A trans-dimensional Bayesian Markov chain Monte Carlo algorithm for model assessment using frequency-domain electromagnetic data. *Geophysical Journal International*, 187(1), 252–272.
- Minsley, B. J., Foks, N. L., & Bedrosian, P. A. (2021). Quantifying model structural uncertainty using airborne electromagnetic data. *Geophysical Journal International*, 224(1), 590–607.
- Minsley, B. J., Rigby, J. R., James, S. R., Burton, B. L., Knierim, K. J., Pace,
   M. D., ... Kress, W. H. (2021). Airborne geophysical surveys of the lower
   mississippi valley demonstrate system-scale mapping of subsurface architecture.
   *Communications Earth & Environment*, 2(1), 1–14.
  - Moghadas, D. (2020). One-dimensional deep learning inversion of electromagnetic induction data using convolutional neural network. *Geophysical Journal International*, 222(1), 247–259.
- Moghadas, D., Behroozmand, A. A., & Christiansen, A. V. (2020). Soil electrical conductivity imaging using a neural network-based forward solver: Applied to large-scale bayesian electromagnetic inversion. Journal of Applied Geophysics, 104012.
- Mosegaard, K., & Tarantola, A. (1995). Monte Carlo sampling of solutions to inverse problems. J. geophys. Res, 100(B7), 12431–12447.
- Mosser, L., Dubrule, O., & Blunt, M. J. (2017). Reconstruction of three-dimensional porous media using generative adversarial neural networks. *Physical Review E*, 960 96(4), 043309.
- Mosser, L., Dubrule, O., & Blunt, M. J. (2020). Stochastic seismic waveform inver sion using generative adversarial networks as a geological prior. Mathematical
   *Geosciences*, 52(1), 53–79.
- Puzyrev, V., & Swidinsky, A. (2019). Inversion of 1d frequency-and time-domain electromagnetic data with convolutional neural networks. *arXiv preprint arXiv:1912.00612*.
- Rimstad, K., Avseth, P., & Omre, H. (2012). Hierarchical Bayesian lithology/fluid
   prediction: A North Sea case study. *Geophysics*, 77(2), B69–B85.
- <sup>969</sup> Röth, G., & Tarantola, A. (1994). Neural networks and inversion of seismic data.
   <sup>970</sup> Journal of Geophysical Research: Solid Earth, 99(B4), 6753–6768.
- Sambridge, M., Bodin, T., Gallagher, K., & Tkalčić, H. (2013). Transdimensional inference in the geosciences. *Philosophical Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences*, 371(1984), 20110547.
- Scales, J. A., & Snieder, R. (1997). To [b]ayes or not to [b]ayes? *Geophysics*, 62(4), 1045–1046.
- Scheidt, C., Renard, P., & Caers, J. (2015). Prediction-focused subsurface modeling:

977	investigating the need for accuracy in flow-based inverse modeling. Mathemati-
978	cal Geosciences, $47(2)$ , 173–191.
979	Shahraeeni, M. S., & Curtis, A. (2011). Fast probabilistic nonlinear petrophysical in-
980	version. $Geophysics$ , $76(2)$ , E45–E58.
981	Smith, B. D., Abraham, J. D., Cannia, J. C., Minsley, B. J., Deszcz-Pan, M., & Ball,
982	L. B. (2010). Helicopter electromagnetic and magnetic geophysical survey
983	data, portions of the North Platte and South Platte Natural Resources Dis-
984	tricts, Western Nebraska (Tech. Rep. No. 2010-1259). U.S. Geological Survey
985	Scientific Investigations Report. doi: 10.3133/ofr20101259
986	Tarantola, A. (2005). Inverse problem theory and methods for model parameter esti-
987	mation. SIAM.
988	Tarantola, A., & Valette, B. (1982a). Generalized nonlinear inverse problems solved
989	using the least squares criterion. Rev. Geophys. Space Phys, $20(2)$ , 219–232.
990	Tarantola, A., & Valette, B. (1982b). Inverse problems= quest for information. $J$ .
991	Geophys, 50(3), 150-170.
992	Tikhonov, A. N. (1963). On the solution of ill-posed problems and the method of
993	regularization. In Doklady akademii nauk (Vol. 151, pp. 501–504).
994	Viezzoli, A., Christiansen, A. V., Auken, E., & Sørensen, K. (2008). Quasi-3D mod-
995	eling of airborne TEM data by spatially constrained inversion. <i>Geophysics</i> ,
996	73(3),  m F105-F113.
997	Vilhelmsen, T. N., Auken, E., Christiansen, A. V., Barfod, A. S., Marker, P. A.,
998	& Bauer-Gottwein, P. (2019). Combining clustering methods with mps to
999	estimate structural uncertainty for hydrological models. Frontiers in Earth
1000	Science, 7, 181.
1001	Zhang, X., & Curtis, A. (2020a). Seismic tomography using variational in-
1002	ference methods. Journal of Geophysical Research: Solid Earth, 125(4),
1003	e2019JB018589.
1004	Zhang, X., & Curtis, A. (2020b). Variational full-waveform inversion. <i>Geophysical</i>
1005	Journal International, 222(1), 406–411.
1006	Zhang, X., & Curtis, A. (2021). Bayesian geophysical inversion using invertible
1007	neural networks. Journal of Geophysical Research: Solid Earth, 126(7),
1008	e2021JB022320.
1009	Zhao, X., Curtis, A., & Zhang, X. (2022). Bayesian seismic tomography using nor-
1010	malizing flows. <i>Geophysical Journal International</i> , 228(1), 213–239.