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Optimal experiment design for bottom friction parameter estimation

Simon C. Warder^{a,*}, Matthew D. Piggott^a

^a*Department of Earth Science and Engineering, Imperial College London, London, SW7 2AZ, UK*

Abstract

It is common practice within numerical coastal ocean modelling to perform model calibration with respect to a bottom friction parameter. While many modelling studies employ a spatially uniform coefficient, within the parameter estimation literature the coefficient is typically taken to be spatially (or even temporally) varying. A parameter estimation experiment requires an appropriate set of observations, and also the selection of an appropriate parameter space which captures the spatial variability of the bottom friction parameter. In regions such as the Bristol Channel, which is used as a case study within this work, observation data is relatively abundant; here we use observations of M2 and S2 harmonic amplitudes and phases at 20 locations within the Channel. However, as is typical within friction parameter estimation problems, there is no obvious constraint on the spatial variation of the friction coefficient. Here, we define the parameter estimation ‘experiment design’ as the mapping from a small number of friction parameters onto the model domain. We propose a robust method for the appropriate selection of a low-dimensional experiment design, utilising an optimal experiment design (OED) technique via construction of the Fisher Information Matrix. The objective is to identify the experiment design resulting in the tightest possible constraints on the unknown parameters, given the available observation data. We construct the Fisher Information Matrix via the use of an adjoint shallow water numerical model, *Thetis*, and perform a variant of D-optimal design to find the optimal experiment design from within two *a priori* choices of design space. These are based on splitting the model domain either by simple slices across the channel, or by the type of sediment found on the sea bed. We first validate the OED framework by utilising a Bayesian inference algorithm to perform parameter estimation using a selection of experiment designs, which confirms that the OED framework offers a good estimate of the true parameter uncertainty resulting from the use of a given experiment design. An exploration of the full space of experiment designs shows that up to three Manning’s n coefficient values can be estimated from the observation data to within an uncertainty of approximately $0.001 \text{ s m}^{-1/3}$, but that the experiment design is highly influential in achieving this threshold, thus demonstrating the value of our approach as a preliminary step in a parameter estimation study. We also investigate the sensitivity of the achievable parameter uncertainty to the availability of observation data and its measurement uncertainty, providing insights useful to the design of future observation surveys. Finally, we further demonstrate our OED framework with an application to a model of the northwest European continental shelf.

Keywords: Optimal experiment design, Fisher Information Matrix, Parameter estimation, Bottom friction, Manning coefficient

*Corresponding author

Email addresses: s.warder15@imperial.ac.uk (Simon C. Warder), m.d.piggott@imperial.ac.uk (Matthew D. Piggott)

1. Introduction

Numerical tidal models are utilised within a wide range of application areas, including sediment and pollutant transport (e.g. Periañez et al. (2013), Chen and Liu (2017), Li et al. (2018)), ecosystems and fisheries (e.g. Marshall et al. (2017), Whomersley et al. (2018)) and marine renewable energy (e.g. Adcock et al. (2015), Neill et al. (2018), Mackie et al. (2020a), Wang and Yang (2020)), as well as underpinning the modelling of coastal hazards including storm surges (Flather, 2000). Accurate tidal models are therefore of significant value, yet any model output is subject to a variety of sources of uncertainty.

Within coastal ocean modelling, a significant source of uncertainty is the parameterisation of bottom friction. Friction between the ocean and the sea bed arises due to a boundary layer at the interface, and due to form drag induced by bathymetry undulations. The resulting momentum loss is not explicitly captured in coastal ocean models, but is instead typically incorporated via a parameterised drag term in the governing equations, which relates the water velocity to a frictional force. There are many possible choices of drag parameterisation (Zhang et al., 2011, Döös et al., 2004), each of which introduces a drag parameter. Furthermore, this parameter in principle varies spatially, due to variations in sea bed roughness and hydrodynamic conditions, as well as potential dependence on model mesh resolution given that the parameterisation represents subgrid-scale processes. Even when it can be related to land or sea floor classification using well-established tables, this parameter still carries significant uncertainty (Mayo et al., 2014). The bottom friction coefficient may also vary over time, e.g. due to morphological changes at the sea bed (Davies and Robins, 2017) or seasonal variations in hydrological conditions (Huybrechts et al., 2021). However, since the coefficient's value cannot be directly measured in the field, this parameter is a source of significant uncertainty within coastal ocean models.

For this reason, it is common within the coastal ocean modelling literature to perform model calibration with respect to the bottom friction coefficient, whereby its value is inferred from observations of hydrodynamic variables. Commonly used observations for friction parameter estimation include timeseries or harmonic analysis data from tide gauges, current measurements (e.g. from ADCPs), or satellite altimetry. Techniques for performing this calibration include Kalman filters (e.g. Mayo et al. (2014)), Bayesian inference approaches (e.g. Hall et al. (2011), Sraj et al. (2014b)), and gradient-based methods utilising adjoint models (e.g. Maßmann (2010)) or other gradient calculation techniques (e.g. Sraj et al. (2014a)). Adjoint-based methods can be efficiently applied to the estimation of very large numbers of control parameters. However, this leads to the problem of overfitting, which is commonly addressed in the literature by either regularisation of the calibration problem (Maßmann, 2010), the reduction of the dimension of the parameter space (Lu and Zhang, 2006, Zhang et al., 2011), truncation of the optimisation after a small number of iterations (Heemink et al., 2002), or combinations of the above. It should be noted that any calibration with respect to the bottom friction parameter will simultaneously be addressing both unknown physical quantities (e.g. related to bottom roughness), as well as errors associated with other model inputs, and modelling and discretisation choices. However, calibration with respect to bottom friction alone is commonly undertaken (e.g. Ullman and Wilson (1998) and several of the above references), and is therefore the focus of this study. The methodology presented can also be applied to calibration with respect to other model inputs.

Any parameter estimation problem requires two components: firstly, a suitable set of observations from which to infer the unknown model parameters, and secondly, an appropriate parameter space within which the unknown

39 parameters lie (e.g. which/how many model parameters to estimate, their spatial distribution, etc). These two aspects
40 constitute the ‘experiment design’, the optimisation of which is known as ‘optimal experiment design’ (OED), and
41 is the topic of this paper. The optimisation of these two aspects is often considered independently.

42 The optimisation of the observation strategy amounts to selecting which variables to measure, and when and
43 where to measure them, in order to infer a given set of unknown model parameters. This aspect of OED is common
44 in the literature, with example applications in a variety of fields (Ucinski, 2004). The computation of the model’s
45 sensitivities with respect to its unknown inputs is central to the OED method, typically via construction of the Fisher
46 Information Matrix (FIM) (Fedorov, 1972) or something closely related. For a given experiment design, the inverse
47 of the FIM gives a lower bound on the covariance matrix of the estimated parameters (Söderström and Stoica, 1989,
48 Machado et al., 2009, Alaña and Theodoropoulos, 2011). The optimisation of the experiment design can be defined
49 as the minimisation of this parameter covariance, so the crux of FIM-based OED methods is therefore to maximise (a
50 scalar measure of) the FIM. Since the FIM is constructed from model sensitivities, it does not explicitly depend on the
51 values of the model inputs or the observations themselves. The OED process can therefore be performed in advance of
52 expensive observation surveys. FIM-based OED methods have been applied in a variety of experimental sciences (e.g.
53 Balsa-Canto et al. (2008), Strigul et al. (2009), Yu et al. (2018)). Within the field of numerical coastal and estuarine
54 modelling, de Brauwere et al. (2009) utilised an FIM-like method for the calibration of a reactive transport model,
55 designing a sampling strategy to best constrain mortality and sedimentation rates of *E. coli* in the Scheldt Estuary
56 using a given number of water samples. Vandenberghe et al. (2002) utilised a similar FIM-based method to design
57 an optimal sampling strategy in a water quality model of the River Dender. Graham et al. (2017) take an alternative
58 sensitivity-based approach to optimising experiment design for a storm surge case study, in order to determine an
59 optimal configuration of observation buoys with which to infer a given set of unknown friction parameters. A similar
60 concept to optimal experiment design is that of ‘targeted observations’ for state estimation (e.g. Langland (2005)).
61 However, while FIM methods have been applied to state estimation (Alaña and Theodoropoulos, 2011), the focus of
62 the present study is on parameter estimation.

63 The second aspect of experiment design is the selection of the parameters to be estimated. This aspect is the
64 focus of the present study since, as is the case in many similar coastal regions, a reasonable set of observation data
65 already exists within the selected study region, while there is no obvious choice for the input parameter space. A
66 similar and commonly approached problem in the OED literature is the selection of a subset of model parameters
67 for estimation; this has often been solved via FIM methods (Machado et al., 2009, Chu and Hahn, 2009, Kravaris
68 et al., 2013) or other sensitivity analysis-based methods (Li et al., 2004, Wang et al., 2018).

69 The concept of experiment design for bottom friction parameter estimation has been discussed in numerous
70 studies within the coastal ocean modelling literature. A common approach to designing such experiments is to
71 select a low-dimensional representation of the bottom friction field, so that only a relatively small number of control
72 parameters need to be inferred from the observation data. However, appropriately selecting the control parameters
73 is not a straightforward task. Various methods have been used for the selection of these control parameters (Lardner
74 et al., 1993, Ullman and Wilson, 1998, Chen et al., 2014), with typical schemes involving distributing control points
75 across the model domain either uniformly, based on some physical quantity such as the bathymetry gradient (Lu
76 and Zhang, 2006), or based on the volume and distribution of the observation data (Das and Lardner, 1991, Altaf

77 et al., 2012). Other studies have divided their model domain into subdomains based on sedimentology data (Guillou
78 and Thiébot, 2016), land classification (Graham et al., 2017) or bathymetry contours (Sraj et al., 2014b). However,
79 due to the difficult nature of the parameter selection problem, most of the above examples still required some form
80 of regularisation in order to avoid overfitting, despite the modest number of control parameters. This regularisation
81 typically takes the form of a penalty term in the misfit functional, which provides a ‘prior’ value for the parameter,
82 thus constraining its value where the observations are insufficient to do so. However, while this constitutes a practical
83 approach, one drawback of such regularisation is that the penalty term introduces additional degrees of freedom (e.g.
84 the magnitude of the penalty term, the prior value for the parameter, etc) which must be selected appropriately, e.g.
85 via a cross-validation approach utilising a second observation dataset (Ullman and Wilson, 1998).

86 Heemink et al. (2002) used an adjoint model to compute the gradient of the model-observation misfit with respect
87 to the spatially varying inputs, in order to inform the selection of a small set of independent control parameters.
88 However, this approach has the disadvantage that the gradient calculation depends on the observation data. As
89 described above, an advantage of OED frameworks is that they are independent of the observation data itself,
90 depending only on the experiment design (i.e. the location and timing of the observations). This enables OED
91 frameworks to be applied in advance of (expensive) observation surveys. This dependence on the observation data
92 also means that the resulting experiment design is implicitly motivated by reducing the model-observation misfit,
93 rather than identifying designs resulting in tight constraints on the control parameters. Furthermore, the approach
94 by Heemink et al. (2002) still did not entirely avoid the overfitting problem, instead relying on truncation of the
95 optimisation algorithm after a small number of iterations. This again requires an additional dataset in order to
96 determine when to terminate the optimisation algorithm, therefore sharing one of the drawbacks of the conventional
97 penalty term regularisation approach.

98 Within this study, we demonstrate the application of an OED framework to a bottom friction parameter estimation
99 problem. We seek a low-dimensional piecewise-constant representation of the spatially varying bottom friction
100 coefficient within a model of the Bristol Channel, in which all selected control parameters can be constrained by the
101 observations, without the need for regularisation. We proceed via an FIM-based method which utilises an adjoint-
102 capable shallow water solver to compute model sensitivities. We use a modified version of the so-called D-criterion
103 as a scalar measure of the FIM. The maximisation of this criterion corresponds to the minimisation of the parameter
104 estimate confidence region, and hence to the experiment design providing the tightest constraints on the unknown
105 parameters. In this respect, the optimal experiment designs identified within this study are motivated by parameter
106 identifiability, and our methodology is therefore distinct from one which is motivated only by minimising model-
107 observation discrepancy, where parameter uncertainty (and consequently model outputs at uncalibrated locations)
108 may remain high. Note also that this approach can be considered as an alternative to the traditional regularisation
109 approach to avoiding under-constrained problems, since it is possible to obtain a set of control parameters which are
110 constrained by the observations (to a given precision), without the need for regularisation or optimisation truncation
111 as described above. To the authors’ knowledge, this paper is the first application of an OED procedure to a coastal
112 ocean parameter estimation problem. Our approach is also novel in its use of the adjoint model for the construction
113 of the FIM, for the purpose of selecting a low-dimensional parameterisation of a fully spatially-varying model input.

114 The adjoint-capable model used within this work, *Thetis*, is described in section 2 along with the case study

115 region and the observation data used. The optimal experiment design problem is described in section 3, including
 116 detail of the FIM-based optimisation method we employ. Results from the OED framework are presented in section
 117 4, including verification that our modified D-criterion provides a reasonably tight bound on the size of the true
 118 parameter confidence region. This section also explores the influence of the quantity and precision of the observations
 119 on experiment design performance. The results from model calibration using our optimal design are summarised in
 120 section 5. In section 6, we further demonstrate the OED framework via an application to a model of the northwest
 121 European continental shelf. The implications of our results are discussed in section 7, including avenues for future
 122 work. Finally, a summary and conclusions can be found in section 8.

123 2. Model and study region

124 Within this work, we take the Bristol Channel and Severn Estuary as a case study region. The Bristol Channel
 125 is a macrotidal inlet situated toward the south-east of the UK, and is of significant interest for its tidal range energy
 126 resource (Angeloudis and Falconer, 2017). It is also the site of the Hinkley Point nuclear power station. Furthermore,
 127 accurate models are vital for flood risk studies (e.g. Lyddon et al. (2018)). As such, the Bristol Channel is a site of
 128 particular interest for the development of accurate numerical models.

129 2.1. Two-dimensional adjoint-capable shallow water model

130 *Thetis* is a 3D (Kärnä et al., 2018, Pan et al., 2020) and 2D (Vouriot et al., 2019) finite element based coastal
 131 ocean model, built using the finite element code generation framework *Firedrake* (Rathgeber et al., 2016). Within
 132 this work, we use *Thetis* in its two-dimensional depth-averaged mode, solving the nonlinear shallow water equations
 133 given by

$$\begin{aligned} \frac{\partial \eta}{\partial t} + \nabla \cdot (H\mathbf{u}) &= 0, \\ \frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} + \mathbf{F}_C + g\nabla \eta &= -\frac{\boldsymbol{\tau}_b}{\rho H} + \nabla \cdot (\nu(\nabla \mathbf{u} + \nabla \mathbf{u}^T)), \end{aligned} \quad (1)$$

134 where η is the surface elevation, \mathbf{u} is the depth-averaged velocity vector, $H = \eta + h$ is the total water depth, h is
 135 the bathymetry (measured positive downwards), \mathbf{F}_C is the Coriolis force, g is the acceleration due to gravity, ρ is
 136 the density, ν is the viscosity, and $\boldsymbol{\tau}_b$ is the bottom stress. Within this work, the bottom stress is parameterised via
 137 Manning's n formulation

$$\frac{\boldsymbol{\tau}_b}{\rho} = \frac{gn^2}{H^{\frac{1}{3}}} |\mathbf{u}| \mathbf{u}, \quad (2)$$

138 where n is the (spatially varying) Manning coefficient. This coefficient is assigned a uniform value of $n = 0.025$
 139 $\text{sm}^{-1/3}$ outside the Bristol Channel, while its value inside the Channel is to be inferred via parameter estimation
 140 methods; see figure 1. Wetting and drying of intertidal regions is handled via the bathymetry modification scheme
 141 of Kärnä et al. (2011). This scheme introduces a wetting-drying parameter α , which controls the transition from wet
 142 to dry regions of the domain. Smaller values of α produce more accurate results, but there exists a minimum stable
 143 value which depends on the bathymetry gradient and mesh resolution at the wet-dry interface. In this work, α is
 144 taken as 0.2 m, which was found by experimentation to be close to the minimum stable value for this model setup.

145 The unstructured mesh used throughout this study is shown in figure 1, and is based on that used in Mackie et al.
 146 (2020b). The mesh was generated on a UTM30 coordinate projection, using the Python package *qmesh* (version

147 1.0.2) (Avdis et al., 2018), which interfaces the mesh generator *Gmsh* (version 3.0.4) (Geuzaine and Remacle, 2009).
 148 The coastline data is from the Global Self-consistent, Hierarchical, High-resolution Geography Database (GSHHG)
 149 (Wessel and Smith, 1996). The mesh resolution varies between 250 m at the innermost part of the Channel, and
 150 8 km at the ocean boundaries, with an intermediate resolution along other parts of the coastline. This results
 151 in a total of 42,862 triangular elements. The bathymetry is taken from Digimap (Digimap, 2013), which has a
 152 resolution of approximately 180 m. The ocean boundaries are forced with tidal elevations generated based on the
 153 dominant constituents (M2 and S2) from the TPXO database (Egbert and Erofeeva, 2002). The next most significant
 154 constituent (N2) has less than half the amplitude of these primary constituents. While the choice to force with
 155 only two constituents neglects any transfer of energy from other tidal constituents, we assume that this forcing is
 156 sufficient for the purposes of this study, which represents a proof of concept for our OED framework. A no-normal-flow
 157 boundary condition is applied on the coastal boundaries, and river outflow is neglected.

158 The governing equations (1) are solved using the $P_1^{\text{DG}}-P_1^{\text{DG}}$ finite element pair, with a Crank-Nicolson timestep-
 159 ping method and a timestep size of $\Delta t = 200$ s. All model runs within this work follow a generic spin-up period of
 160 10 days with a default uniform friction parameter of $n = 0.025 \text{ s m}^{-1/3}$. Following this initial spin-up, each friction
 161 configuration is run for a further 17.77 days, with harmonic analysis (for the M2 and S2 constituents) performed on
 162 the final 14.77 days. This period was chosen to meet the Rayleigh criterion for resolving the M2 and S2 constituents.

163 The adjoint mode of Thetis, which is used within this work for the computation of model sensitivity with respect
 164 to the Manning coefficient n , is generated via *pyadjoint* (Mitusch et al., 2019, Farrell et al., 2013), which interfaces
 165 with Firedrake to efficiently derive the adjoint model. For more detail on Thetis adjoint and its applications, see
 166 previous studies e.g. Warder et al. (2019, 2021), Goss et al. (2020).

167 2.2. Observation data

168 The observations used within this study are the M2 and S2 harmonic amplitudes and phases at the 20 tide gauge
 169 locations indicated in figure 1 (National Oceanography Centre, personal communication 2018). This amounts to
 170 a total of 80 pieces of data, which can be compared with model outputs. While more harmonic constituents are
 171 available within the dataset, the M2 and S2 constituents dominate the dynamics within the Bristol Channel with
 172 typical amplitudes of around 4 m and 1.5 m respectively, compared with around 0.6 m for the next most significant
 173 constituent (N2). The selected OED methodology incurs a computational cost proportional to the volume of data
 174 used for the parameter estimation, so we choose not to consider additional constituents and focus only on the M2
 175 and S2 harmonics. The effect of the volume of observation data (including the choice of constituents) on experiment
 176 design performance is considered in section 4.3.

177 Throughout this work, we assume that some uncertainty is introduced by the observations. Specifically, we
 178 assume that the amplitude observations are independent and identically distributed variables, and similarly for the
 179 phases, with variances of 25 cm^2 and 6.25°^2 , respectively. While this is somewhat arbitrary, we discuss the influence
 180 of this choice in section 4.4.

181 3. Optimal experiment design

182 We aim to determine an optimal low-dimensional representation of the spatially varying Manning coefficient field.
 183 In particular, we seek a piecewise-constant field, such that the coefficient field is specified by a vector of length m .

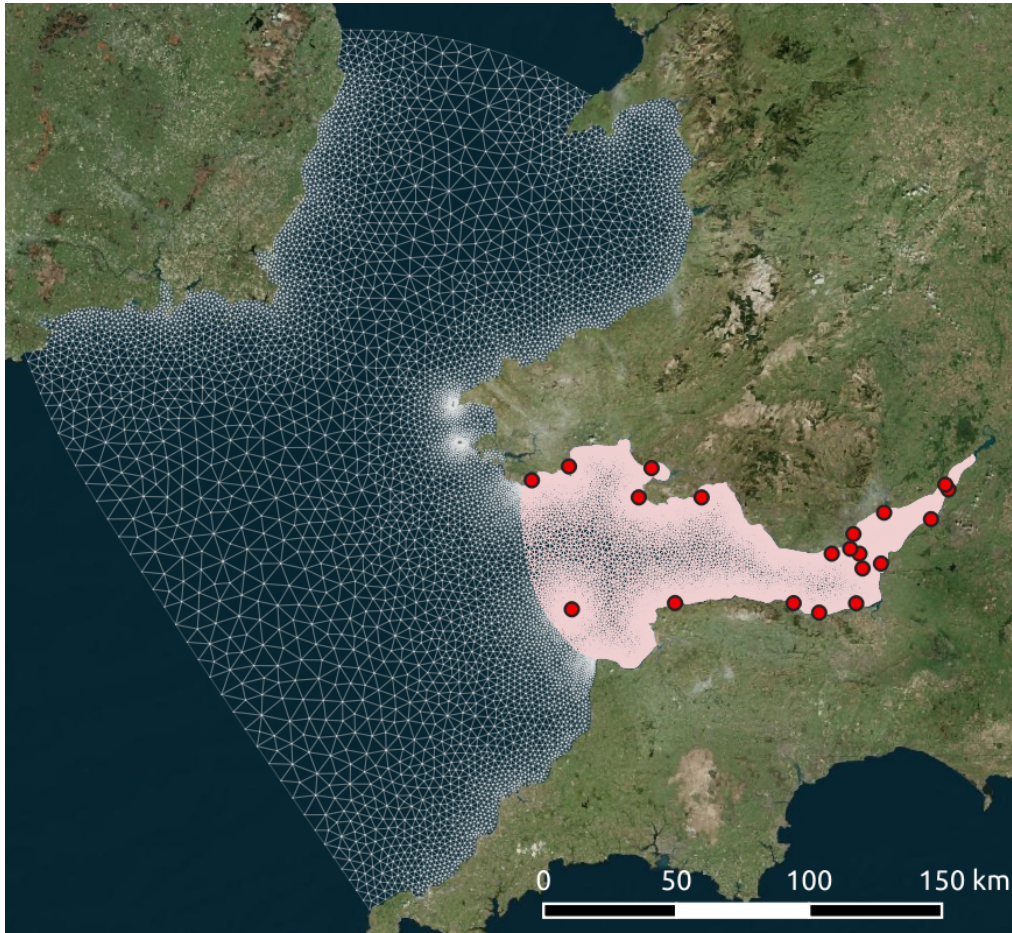


Figure 1: Unstructured mesh used for all simulations within this work. Pink indicates the region where the friction coefficient is permitted to vary spatially. Red circles indicate locations where tidal harmonic data is used.

184 The aim is to determine the optimal number of parameters m , and the optimal mapping from these parameters
 185 onto the model domain, such that tight constraints on the parameters can be achieved via a parameter estimation
 186 exercise using the observations described in section 2.2. The choice of this mapping and number of parameters will
 187 be referred to as the experiment design. We emphasise that this problem is distinct from conventional applications of
 188 OED to parameter selection, where an experiment design typically consists of a selection of a subset of parameters,
 189 with unselected parameters held constant (e.g. Li et al. (2004), Chu and Hahn (2007), Machado et al. (2009)).

190 We first apply an *a priori* constraint on the friction coefficient, by splitting the model domain into subdomains,
 191 within each of which the friction coefficient is assumed to be uniform. This constraint permits the use of an exhaustive
 192 search algorithm to find the optimal design. The selected number of these subdomains must be sufficiently large that
 193 the space of experiment designs contains a ‘good’ design, but small enough to facilitate the use of the exhaustive
 194 search algorithm. Within this work, we consider two such *a priori* constraints, which are described in section 3.1.

195 Once a suitable subdomain parameterisation of the friction coefficient has been selected, we explore the remaining
 196 space of experiment designs to determine the optimal design. Each design corresponds to a grouping of the *a priori*
 197 subdomains into m groups, such that the friction coefficient is specified by m parameters. In order to rank the
 198 experiment designs, we compute the Fisher Information Matrix corresponding to each design, and find the design

199 which optimises a selected scalar measure of this matrix. This approach is described in detail in section 3.2.

200 3.1. *A priori* subdomain parameterisations

201 Within this work, we consider two different choices for the *a priori* subdomain parameterisation of the friction
202 coefficient field:

203 1. Division of the domain by slices drawn by hand, as shown in figure 2. This division is somewhat arbitrary;
204 the number of subdomains was chosen to be the same as the number of sediment classification types present in
205 the domain (eight, see below), with the slices chosen to divide the Bristol Channel fairly evenly, with slightly
206 smaller sections further into the estuary where the density of observation locations is higher. The possible
207 experiment designs are further constrained such that the slices are grouped into contiguous blocks (i.e. region
208 1 and region 3 cannot be assigned the same friction parameter unless region 2 is also included).

209 2. Division of the domain based on sediment data, as shown in figure 3. The sediment data is taken from the SHOM
210 database (Service Hydrographique et Océanographique de la Marine) (SHOM, 2013), which divides the domain
211 into discrete sediment types. Since the presence of sediment influences the physical process of momentum loss
212 due to bottom friction, this is an attempt to incorporate physical knowledge into the experiment design. The
213 sediment types within the model domain are shown in figure 3. The candidate experiment designs are further
214 constrained by requiring that the sediment-based subdomains are grouped contiguously in order of sediment
215 grain size; this order is shown in table 1, along with the area occupied by each sediment type within the model
216 domain.

217 Each possible experiment design corresponds to a grouping of the selected subdomains, according to the *a priori*
218 constraints described above. Given that there are eight subdomains, for a parameterisation with $m = 2$ (i.e. a
219 grouping of the subdomains into two groups), there are seven possible designs for each choice of *a priori* constraint.
220 For $m = 3$, there are 21 possible designs, and for $m = 4$ there are 35.

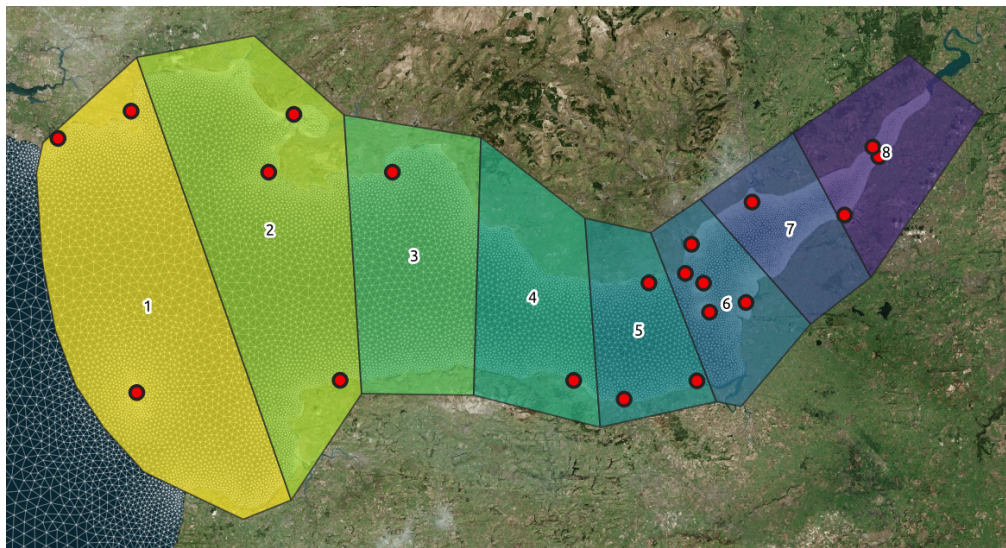


Figure 2: Division of the Bristol Channel into slices, labelled by ID number. Red circles indicate locations where tidal harmonic data is used.

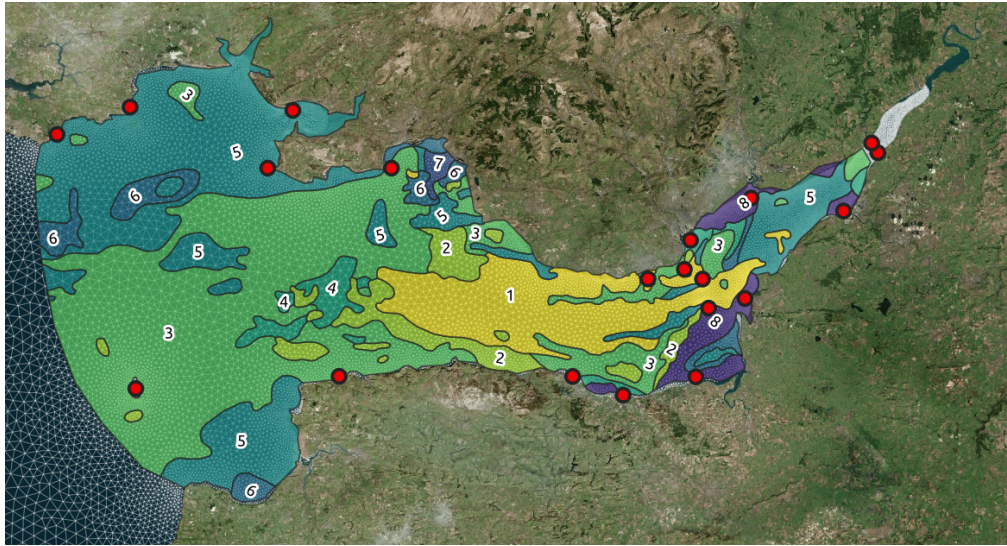


Figure 3: Spatial distribution of sediment types within the Bristol Channel, labelled by ID number. Red circles indicate locations where tidal harmonic data is used.

Sediment name	NFRoche	NFG	NFSG	NFSGV	NFS	NFSV	NFVS	NFV
ID	1	2	3	4	5	6	7	8
Area covered / km ²	761	375	2800	199	2051	297	28	216

Table 1: Sediment types found within the Bristol Channel (from SHOM dataset), sorted by roughness length.

221 3.2. The Fisher Information Matrix

222 For reviews of parameter selection methods via the Fisher Information Matrix (FIM), the reader is referred to
 223 works from other fields, e.g. Machado et al. (2009), Kravaris et al. (2013). However, note that the problem considered
 224 here is subtly different to the conventional parameter selection problem, where the objective is typically to select an
 225 optimal subset from a large set of parameters, with unselected parameters taking prescribed values. In this work, we
 226 seek an experiment design in which the friction parameter is specified everywhere within the Bristol Channel; each
 227 experiment design corresponds to a *grouping* of model subdomains, not a subset selection. This is perhaps more
 228 similar to the clustering of parameters considered by Chu and Hahn (2009), although we take a different approach.
 229 The following is a simple exposition of the FIM approach to optimal experiment design.

230 Consider a parameter estimation problem, where a vector of unknown parameters $\boldsymbol{\theta}$ is to be estimated by min-
 231 imising a misfit functional given by

$$(\mathbf{y} - \mathbf{f}(\boldsymbol{\theta}))^T \text{cov}_e^{-1} (\mathbf{y} - \mathbf{f}(\boldsymbol{\theta})), \quad (3)$$

232 with respect to $\boldsymbol{\theta}$, where \mathbf{y} and $\mathbf{f}(\boldsymbol{\theta})$ are vectors of observed and modelled quantities respectively, and cov_e is the
 233 observation error covariance matrix. The parameter covariance matrix resulting from the minimisation of equation
 234 (3) can be estimated as

$$\text{cov}_p = (J^T \text{cov}_e^{-1} J)^{-1}, \quad (4)$$

235 where J is the Jacobian matrix with elements given by

$$J_{ij} = \frac{\partial f_i(\boldsymbol{\theta})}{\partial \theta_j}. \quad (5)$$

236 It is common within the literature to utilise a forward difference method to compute this Jacobian matrix. In this
237 work we instead use the adjoint mode of the numerical model; this is described in section 3.2.3.

238 The parameter covariance given by equation (4) does not depend on the observations \mathbf{y} , but only on their error
239 covariance. Note, however, that it implicitly depends on the number, location and timing of the observations, and
240 other aspects of the experiment design. For a given experiment design, the parameter covariance can therefore be
241 estimated based only on the model sensitivities and the observation covariance, and hence can be estimated prior
242 to solving the parameter estimation problem, and without committing resources to observation surveys or model
243 inversion.

244 Within the optimal experiment design literature, it is common to work with the inverse of the parameter covariance
245 matrix, the so-called Fisher Information Matrix (FIM) (Fedorov, 1972), defined as

$$\text{FIM} = J^T \text{cov}_e^{-1} J. \quad (6)$$

246 Optimisation of the experiment design is undertaken by finding the configuration (in general, the timing and location
247 of observations, or selection of the input parameter space) which optimises some scalar measure of the FIM. There
248 are a variety of common choices for this scalar measure in the literature (Emery and Nenarokomov, 1998, Kravaris
249 et al., 2013).

250 Within this work, we use a modified version of the so-called D-criterion, which we motivate here. In the case
251 of estimating a single parameter, the optimal experiment design can be defined as the one which minimises the
252 confidence interval of the estimated parameter. The higher-dimensional analogue of this is that the optimal design
253 minimises the volume of the confidence ellipsoid of the estimated parameters. This volume is characterised by the
254 determinant of the parameter covariance matrix, $\det(\text{cov}_p)$ (Walter and Pronzato, 1990). Since the FIM is the
255 inverse of cov_p , the minimisation of this confidence ellipsoid volume corresponds to the maximisation of $\det(\text{FIM})$.
256 The so-called D-criterion is simply defined as this determinant, $\det(\text{FIM})$.

257 Within this work, we define a modified D-criterion given by

$$\text{ModD} = \det(\text{FIM})^{1/m}, \quad (7)$$

258 where m is the number of parameters to be estimated. This is sometimes given as the standard definition of the D-
259 criterion (e.g. Emery and Nenarokomov (1998)), but for clarity we refer to this as the modified D-criterion, or ModD.
260 In terms of parameter confidence ellipsoids, the maximisation of the ModD criterion corresponds to the minimisation
261 of the geometric mean of the principal dimensions of the parameter confidence ellipsoid, whereas the standard D-
262 criterion corresponds to minimising the parameter confidence ellipsoid volume. For a fixed m , the experiment design
263 which maximises ModD also maximises the standard D-criterion. However, the inverse power of m ensures that the
264 dimensions of the ModD criterion are independent of the number of parameters estimated, so that its value can be
265 directly compared between experiment designs with different numbers of parameters.

266 *3.2.1. A note on nonlinearity*

267 The above exposition is only strictly valid when the model is linear. If the model is nonlinear with respect to
 268 the unknown parameters (as is the case within this work, since the numerical tidal model is nonlinear), equation
 269 (4) is not exact, but the right-hand-side is nevertheless a measure of the maximum achievable parameter precision
 270 (de Brauwere et al., 2009). The inverse of the FIM (as defined by equation (6)) is considered to give a lower bound
 271 on the parameter error covariance matrix (Alaña and Theodoropoulos, 2011). That is,

$$\text{cov}_p \geq \text{FIM}^{-1}, \quad (8)$$

272 where the inequality is understood to mean that $(\text{cov}_p - \text{FIM}^{-1})$ is positive semidefinite (Söderström and Stoica,
 273 1989, Emery and Nenarokomov, 1998, Machado et al., 2009).

274 Nonlinearity also means that the Jacobian, and therefore the FIM, are functions of the unknown parameters.
 275 The optimisation of the modified D-criterion as defined above therefore identifies only a ‘locally optimal’ experiment
 276 design (Ford et al., 1989, Huan and Marzouk, 2013). That is, the optimal experiment design is optimal only within
 277 a region of the parameter space local to the initial parameter estimate, where the model response is approximately
 278 linear. This issue can be overcome with sequential or iterative design (Blanchet et al., 2008, Catania and Paladino,
 279 2009), or maximin optimisation (Pepelyshev et al., 2004, Rojas et al., 2007, Sun, 2007, Ushijima and Yeh, 2015),
 280 but such designs are outside the scope of this study. Instead, it is assumed that an initial estimate for the friction
 281 coefficient of $n = 0.025 \text{ s m}^{-1/3}$ is appropriate, and that the model response to perturbations in the friction coefficient
 282 from this value can be assumed to be linear. The results presented in section 4.1 suggest that this is an acceptable
 283 assumption, and that the bound provided by equation (8) is reasonably tight.

284 *3.2.2. Defining a ‘good’ experiment design*

285 It is useful to consider the values of the ModD criterion for which the corresponding experiment design produces
 286 acceptable constraints on the estimated parameters. This is particularly important when comparing experiment
 287 designs with different numbers of unknown parameters. In the context of estimating Manning coefficients, which
 288 typically vary between around 0.01 and $0.04 \text{ s m}^{-1/3}$ depending on bed composition (Arcement and Schneider, 1989),
 289 a ‘good’ parameter estimate variance might be $10^{-6} \text{ s}^2 \text{ m}^{-2/3}$. Extending this to higher dimensions analogously to
 290 the definition of the ModD criterion, we therefore seek designs which satisfy

$$\det(\text{cov}_p)^{1/m} < 10^{-6} \text{ s}^2 \text{ m}^{-2/3},$$

291 and therefore

$$\det(\text{cov}_p)^{-1/m} > 10^6 \text{ m}^{2/3} \text{ s}^{-2}. \quad (9)$$

292 Since the FIM provides a bound on cov_p via equation (8), we therefore seek designs satisfying

$$\det(\text{FIM})^{1/m} > 10^6 \text{ m}^{2/3} \text{ s}^{-2},$$

293 or equivalently

$$\text{ModD} > 10^6 \text{ m}^{2/3} \text{ s}^{-2}. \quad (10)$$

294 Due to the inequality in equation (8), an experiment design satisfying equation (10) is not guaranteed to achieve
 295 the target parameter estimate precision given by equation (9). However, we show in section 4.1 that the bound

296 provided by the FIM on cov_p appears to be reasonably tight, and therefore that equation (10) is a good indicator of
 297 experiment designs satisfying equation (9). We emphasise that the value of the OED method is that ModD can be
 298 computed in advance of the parameter estimation procedure, and therefore equation (10) provides a useful *a priori*
 299 indicator of experiment design performance.

300 3.2.3. Construction of the Jacobian via the adjoint model

301 The crux of computing the FIM is the construction of the Jacobian (equation (5)), which consists of the sensitivity
 302 of each model output with respect to each unknown parameter. For this application, the model outputs are the M2
 303 and S2 harmonic amplitudes and phases at 20 locations within the domain, and the model inputs correspond to
 304 the set of friction parameters, denoted $\boldsymbol{\theta}$. The difficulty in computing these sensitivities for an arbitrary experiment
 305 design arises from the fact that the mapping from the friction parameters onto the model domain is unique to each
 306 experiment design. Here, we use the numerical adjoint model to overcome this issue.

307 An arbitrary spatially varying Manning coefficient field can be represented by a vector \mathbf{n} of length N , with the
 308 elements of the vector corresponding to the value of the Manning coefficient at each mesh node, and N the number
 309 of mesh nodes within the region of variable friction. For a given experiment design, denoted \mathcal{D} , we have

$$\mathbf{n} = \mathcal{D} \cdot \boldsymbol{\theta}, \quad (11)$$

310 where $\boldsymbol{\theta}$ is the parameter vector of length m . The experiment design \mathcal{D} is an $N \times m$ matrix consisting of ones and
 311 zeros, which maps the parameters $\boldsymbol{\theta}$ onto the mesh nodes.

312 First, we use the adjoint model to compute the Jacobian with respect to \mathbf{n} , given by

$$\frac{\partial f_i}{\partial n_j}. \quad (12)$$

313 This requires one adjoint model run for each observation i , i.e. for each row of the above matrix. There are 80
 314 observations (M2 and S2 amplitude and phase at 20 locations), so the computation of this Jacobian requires 80
 315 adjoint runs. The Jacobian matrix with respect to the parameter vector $\boldsymbol{\theta}$ can then be computed as

$$\begin{aligned} J_{ij} &= \frac{\partial f_i}{\partial \theta_j} \\ &= \frac{\partial f_i}{\partial n_k} \cdot \frac{\partial n_k}{\partial \theta_j} \\ &= \frac{\partial f_i}{\partial n_k} \cdot \mathcal{D}_{kj}, \end{aligned} \quad (13)$$

316 where the $\frac{\partial f_i}{\partial n_k}$ term has been computed via the set of adjoint model runs as described above.

317 Note that, for a given model setup (mesh, etc), the adjoint model runs are a one-off computational overhead,
 318 and the Jacobian with respect to the parameter vector can be computed via equation (13) at negligible additional
 319 computational cost, for an arbitrary design \mathcal{D} . The computational cost of each adjoint model run is on the same
 320 order of magnitude as that of a forward model run. Note that the common approach in the literature to construct
 321 Jacobians for OED is to use forward differences, based on running the forward model with perturbed values for
 322 each model input. While such an approach would be feasible within this study due the relatively strong *a priori*
 323 constraints described in section 3.1, it would rapidly become infeasible as a greater space of possible experiment
 324 designs is considered. In the most general case, the forward differences approach would require $N + 1$ forward model

325 runs, where N is $\mathcal{O}(10,000)$ for the mesh used in this work; this is clearly not computationally feasible. The use of
 326 an adjoint model is a prerequisite for the extension of this OED method to more complex experiment design spaces,
 327 and we therefore take the adjoint approach here. To the authors’ knowledge this is the first OED study to apply
 328 adjoint methods for this purpose.

329 4. OED Results

330 In section 4.1, we verify that the ModD criterion, computed via the FIM, provides a reliable indicator of experiment
 331 design performance. Section 4.2 then summarises the optimal designs returned by the OED framework, as a function
 332 of the number of parameters m and the choice of *a priori* subdomain division. Finally, sections 4.3 and 4.4 address
 333 the effects of data quantity and measurement uncertainty on experiment design performance.

334 4.1. Verification of the ModD criterion as a measure of design performance

335 The effect of the inequality in equation (8), which represents the nonlinearity of the system, is that the FIM
 336 approach is expected to overestimate the true performance of a given experiment design. Here, the aim is to
 337 investigate the significance of this inequality, i.e. to verify whether the FIM approach is suitable for identifying
 338 ‘good’ experiment designs within the friction parameter estimation context. We proceed by utilising a Bayesian
 339 inference framework to solve the parameter estimation problem corresponding to a selection of experiment designs,
 340 computing the resulting parameter covariance matrix for each. We can then compare the ModD criterion with the
 341 equivalent measure of this parameter covariance matrix, given by

$$\det(\text{cov}_p)^{-1/m}. \quad (14)$$

342 If the inequality of equation (8) were replaced with an equality, the expression given by equation (14) would exactly
 343 equal the ModD criterion. By comparing equation (14) with ModD, we can therefore investigate the significance of
 344 the inequality in equation (8) in characterising experiment designs.

345 While the OED approach described in this paper is applicable to a variety of parameter estimation procedures,
 346 here we choose to use a Bayesian inference approach via a Markov Chain Monte Carlo (MCMC) algorithm to solve
 347 selected parameter estimation problems, since it is a straightforward and well-studied method which yields a direct
 348 estimate of the parameter covariance matrix. Details of the MCMC algorithm, which utilises a Gaussian process
 349 emulator as a surrogate for the full numerical model, can be found in Appendix A, and here we simply summarise
 350 the results.

351 We test the experiment designs using both ‘synthetic’ and ‘real’ experiments. In the synthetic cases, we use the
 352 outputs from the numerical model as observation data within the parameter estimation algorithm. A synthetic result
 353 is produced separately for each experiment design, with a friction coefficient field consistent with the design. This
 354 ensures that the optimum model-observation error is exactly zero for each parameter estimation experiment. While
 355 in this case the ‘observation’ uncertainty is also zero, within the Bayesian inference algorithm we specify observation
 356 uncertainties as described in section 2.2, in order to generate parameter covariances consistent with our assumptions
 357 about the real observation uncertainties. The ‘real’ experiments use the real-world observations, where we again
 358 specify the observation uncertainties as described in section 2.2.

359 For a selection of experiment designs, table 2 presents values of both the ModD criterion and equation (14)
 360 resulting from the synthetic and real experiments. We find that, while the $\det(\text{cov}_p)^{-1/m}$ values are smaller than the
 361 ModD values (as expected due to the inequality in equation (8)), they are reasonably close to the bound provided
 362 by ModD, and therefore that the ModD value is a good indicator of experiment design performance.

Experiment design	ModD [$\text{m}^{2/3} \text{s}^{-2}$]	$\det(\text{cov}_p)^{-1/m}$ [$\text{m}^{2/3} \text{s}^{-2}$] using synthetic observations	$\det(\text{cov}_p)^{-1/m}$ [$\text{m}^{2/3} \text{s}^{-2}$] using real observations
Slice-based designs			
Optimal $m = 2$ design	5.98×10^6	5.63×10^6	5.95×10^6
Optimal $m = 3$ design	1.73×10^6	1.25×10^6	1.62×10^6
Worst $m = 3$ design	0.365×10^6	0.357×10^6	0.357×10^6
Optimal $m = 4$ design	0.824×10^6	0.801×10^6	0.794×10^6
Sediment-based designs			
Optimal $m = 3$ design	0.871×10^6	0.691×10^6	0.825×10^6

Table 2: Comparison of parameter estimation experiment performance using synthetic and real data.

363 4.2. Exploration of optimal designs

364 Having verified in section 4.1 that the ModD criterion constitutes a good metric of experiment design performance,
 365 here we explore the full space of possible experiment designs (for $m = 2, 3, 4$ and 5). Figure 4 shows the ModD
 366 criterion values returned by all possible experiment designs, as a function of the number of parameters in the design,
 367 m , for both choices of *a priori* constraint as summarised by figures 2 and 3. The optimal designs for each value of m
 368 are shown in figures 5 and 6 for the slice and sediment subdomains, respectively. We make the following observations:

- 369 (i) The maximum achievable values of the ModD criterion decrease as the number of unknown parameters m
 370 increases. This is expected, since the information provided by the observation data remains constant, but is
 371 used to attempt to constrain an increasing number of parameters.
- 372 (ii) The designs based on slice subdomains perform better than those based on sediment subdomains. This can
 373 be attributed to the fact that several of the sediment types (2, 4, 6, 7, 8) correspond to only small regions
 374 of the model domain. The performance of a given grouping of sediment types is therefore dominated by the
 375 subdomains with the largest area (1, 3, 5). The space of possible experiment designs is effectively smaller than
 376 for the slice-based subdomains, and the maximum achievable ModD criteria are therefore reduced.
- 377 (iii) For the slice subdomains, we find that it is possible to achieve the target criterion ($10^6 \text{ m}^{2/3} \text{ s}^{-2}$) when grouping
 378 the friction subdomains into two or three groups ($m = 2$ or 3), but no four-parameter designs meet the threshold
 379 ModD criterion.
- 380 (iv) For sediment subdomains, we find that it is only possible to meet the threshold ModD criterion using $m = 2$,
 381 with no possible designs for $m \geq 3$ exceeding the threshold.
- 382 (v) In most cases, the ranges of possible ModD criteria for successive values of m overlap. For example, the
 383 best possible $m = 4$ designs outperform the worst possible $m = 3$ designs. This emphasises the value of an

384 OED procedure, since a poor choice of experiment design can be wasteful with the information provided by
 385 the observations, and it may be possible to achieve a tighter constraint on a larger number of parameters by
 386 optimising the experiment design.

387 (vi) For the optimal designs based on slice subdomains (figure 5), the observation locations are fairly equally
 388 distributed between the regions corresponding to each parameter. This is not apparent in the optimal sediment-
 389 based designs (figure 6), due to the groups not being spatially contiguous.

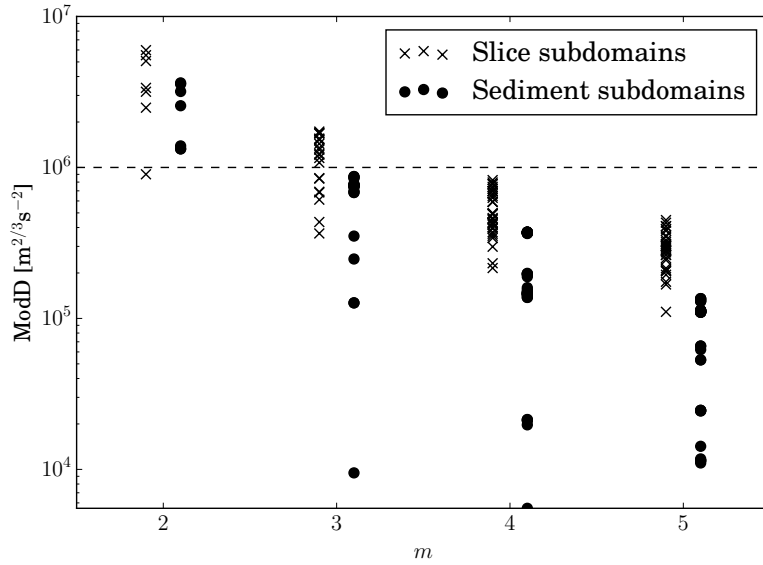


Figure 4: ModD criterion values as a function of the number of parameters in the design, m . Results are shown for both choices of a *priori* subdomain division. The horizontal dashed line indicates the threshold ModD criterion value as described in section 3.2.2.

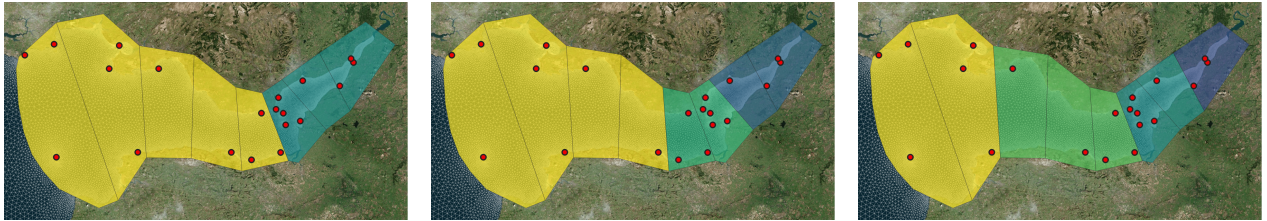


Figure 5: Optimal designs for $m = 2$ (left), $m = 3$ (centre) and $m = 4$ (right), based on the use of slice subdomains.

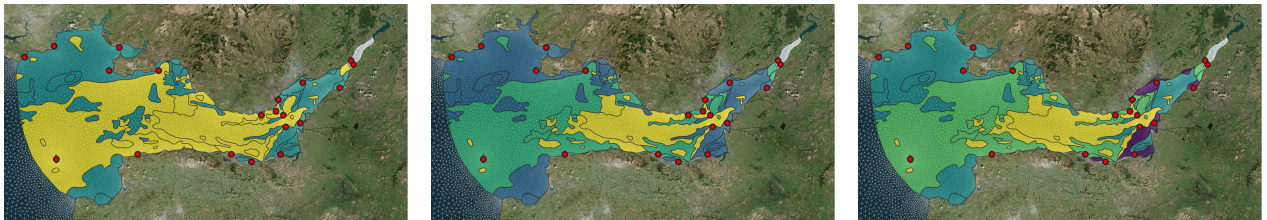


Figure 6: Optimal designs for $m = 2$ (left), $m = 3$ (centre) and $m = 4$ (right), based on the use of sediment subdomains.

4.3. Response of OED criterion to data availability

Here, we investigate the response of the optimal designs, and their corresponding ModD values, to the availability of observation data. To do this, we perform the OED procedure using subsets of the available observations, varying the number of observation locations, and the types of observation data used (i.e. M2/S2 harmonic constituents, amplitude/phase data). This reveals the redundancy in the information provided by the observation dataset, and gives insight into the properties of an experiment design which contribute to its efficacy within parameter estimation.

Figure 7 shows the optimal value of the ModD criterion (on the y -axis), and the corresponding optimal experiment design (indicated by symbols as detailed in the legend), as a function of the number of gauges used, for 3-parameter designs based on slice subdomains using various combinations of observation types. Equivalent results from the use of sediment data for the *a priori* subdomains are shown in figure 8. We make the following observations:

- (i) The sensitivity of the ModD criterion to the number of gauges used decreases, and appears to approach zero, as the number of gauges increases. This can be attributed to the fact that the observation datasets contain redundant information, especially since many of the observation locations are clustered together spatially (see figure 1).
- (ii) For small numbers of observations, the greatest increase in the ModD value is achieved by adding more gauge locations. For large numbers of observations, the ModD value is most efficiently increased by including different observation types (observations of phases in addition to amplitudes, and/or the inclusion of more harmonic constituents in the analysis). This result has possible implications for observation strategies, since it provides guidance on whether it is most efficient to spread observation infrastructure over a wider area (i.e. add more gauges) or focus on more detailed observations in one location (i.e. observe more constituents).
- (iii) For the slice-based subdomains, the optimal design for the maximum number of gauges is the same for all combinations of observation types. Furthermore, only six different optimal designs are found, for any combination of number of gauges and observation types. In other words, the optimal design is relatively insensitive to the volume of data or the types of observations considered. In contrast, the optimal designs are more variable for the sediment-subdomain case. This can be attributed to the fact that many of the sediment types correspond to only very small areas of the Bristol Channel, and subtly different designs therefore have very similar performance, as was shown to be the case in section 4.2.

As a result of the above observations, we conclude that the *a priori* use of slice-based subdomains gives cleaner results, which are easier to interpret, than those based on sediment types, as well as producing overall higher values of the ModD criterion and performing better in parameter estimation problems in terms of the achievable parameter constraints. This is largely due to the more even division of the model domain into slices, compared with the sediment-based division, which resulted in several subdomains of small area. Note, however, that this result does not necessarily suggest that the use of sediment data to constrain the spatial distribution of the friction parameter is detrimental to the model performance, or that the sediment data is inconsistent with the observations.

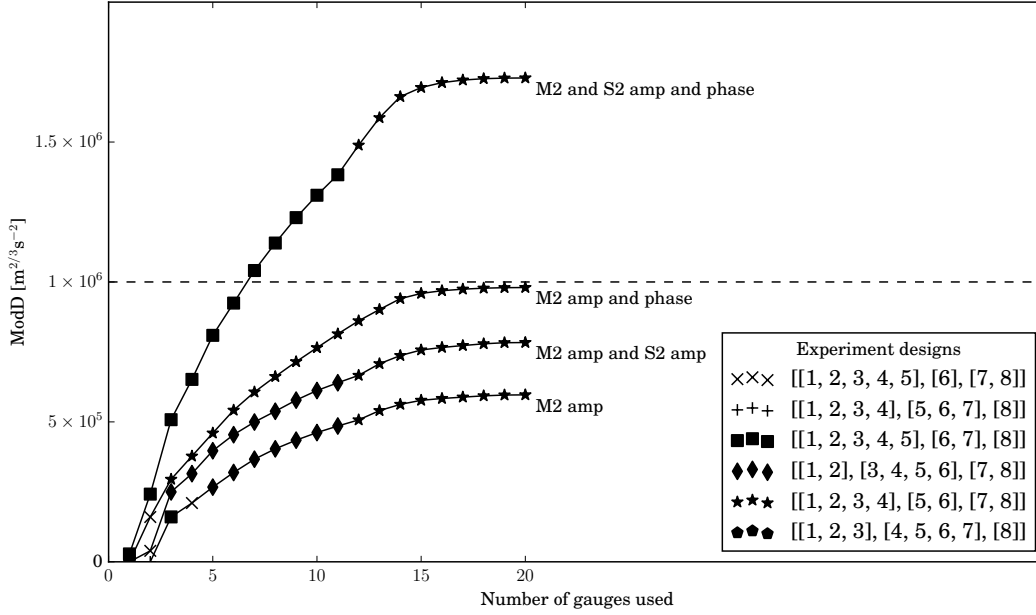


Figure 7: ModD criterion (y -axis) and optimal experiment design (symbols, see legend), as a function of the number of tide gauges used, for four combinations of observation types, based on the use of slice subdomains with $m = 3$. As the number of observation locations increases, the achievable ModD criterion increases. Similarly, the use of multiple types of observations (i.e. phases in addition to amplitudes, or the use of data for multiple harmonic constituents) increases the achievable ModD criterion. The optimal experiment design (i.e. the optimal grouping of friction subdomains) is also dependent on the exact set of gauges and observation types used.

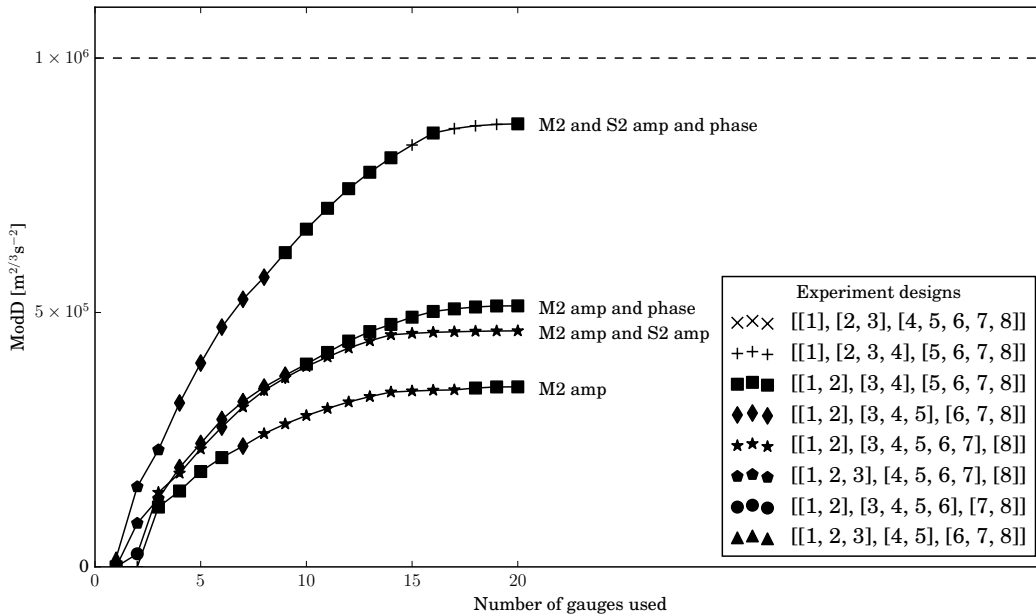


Figure 8: ModD criterion (y -axis) and optimal experiment design (symbols, see legend), as a function of the number of tide gauges used, for four combinations of observation types, based on the use of sediment subdomains with $m = 3$. The results show similar features to those of figure 7.

424 *4.4. Response of OED criterion to observation uncertainty*

425 In this section, we briefly consider the response of the optimal designs to changes in the estimate of cov_e . In
 426 many optimal experiment design studies, the observation covariance matrix cov_e is considered to be proportional to
 427 the identity matrix and is therefore omitted from the formulation of equation (6). However, here we are considering
 428 observations of different dimensions (amplitude and phase), and also compare ModD criterion values with a threshold
 429 value; the inclusion of cov_e is therefore necessary.

430 Equation (6) implies that scaling cov_e by a constant simply scales the ModD values inversely. For example,
 431 decreasing cov_e by a constant factor would have the effect of translating the data of figure 4 in the positive y -
 432 direction, meaning that a greater number of experiment designs would meet the ModD threshold value. In other
 433 words, observations with smaller uncertainty could be used to achieve a given constraint on a larger number of
 434 unknown parameters, or equivalently, fewer observations would be required to achieve a given parameter constraint.

435 In terms of impact on future observation surveys, it is perhaps useful to note the possible trade-off between low
 436 observation uncertainty and the compilation of a diverse set of observations. Based on figure 7, we note that there
 437 is a factor of around three between the ModD values for the assimilation of all data (M2 and S2 amplitude and
 438 phase), compared with M2 amplitude data only. To achieve the same increase in the ModD criterion using only M2
 439 amplitude data would require a reduction by the same factor in the observation error variance. This demonstrates
 440 the balance which can be struck between investing in higher-precision observation techniques and diversifying the
 441 observations made. For example, the OED methodology of this study could be used to compare the value of a large
 442 number of cheap but high-uncertainty observations with a smaller set of expensive but low-uncertainty observations.

443 **5. Performance of the calibrated model**

444 Here we briefly summarise the results from performing the model calibration, using the optimal $m = 3$ slice-based
 445 experiment design. Figure 9 shows the posterior probability density function (PDF) resulting from the MCMC
 446 parameter estimation algorithm, which demonstrates the tight constraints on the control parameters achieved with
 447 this experiment design. To construct the ‘calibrated’ model, we take the mean of this posterior parameter PDF,
 448 resulting in the friction coefficient field depicted in figure 10.

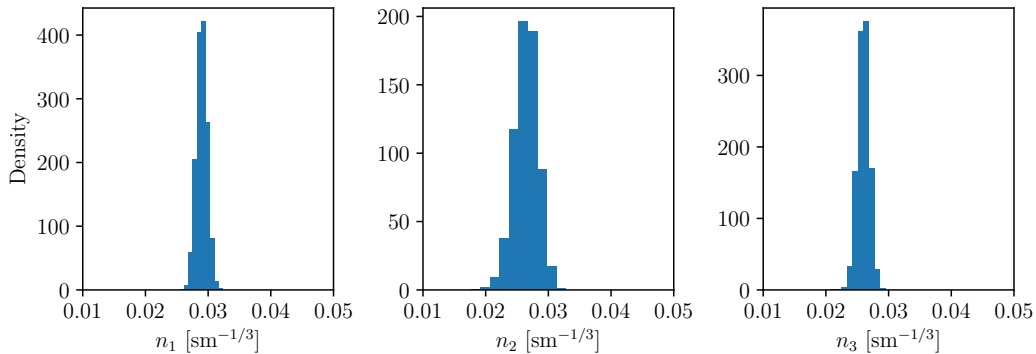


Figure 9: Posterior probability density function for parameter estimation using the optimal $m = 3$ slice-based experiment design.

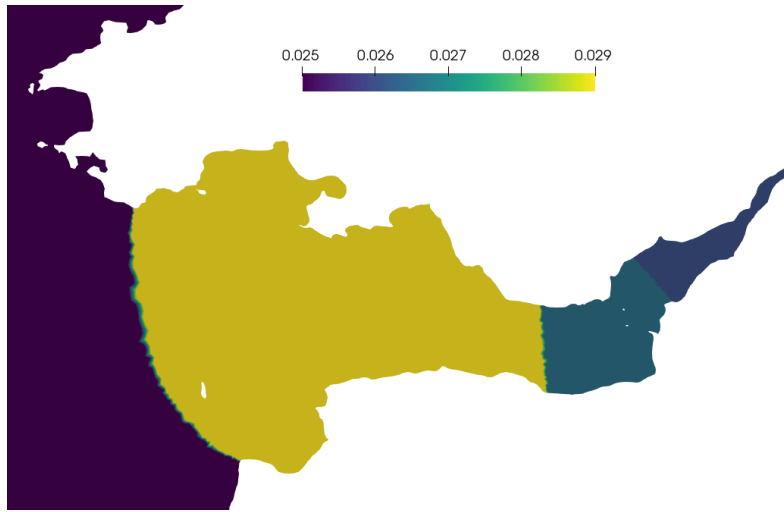


Figure 10: Calibrated Manning coefficient field, using the optimal $m = 3$ slice-based experiment design. Units: $\text{s m}^{-1/3}$.

449 To demonstrate the improvement in performance achieved via this model calibration, we make a comparison with
 450 a uniform parameter with with $n = 0.025 \text{ s m}^{-1/3}$. This comparison is shown in figure 11. The reduction in scatter
 451 between the modelled and observed values is easily visible by eye, and is also evident in the root mean squared
 452 errors, which reduce from 9.7 cm and 4.1° before calibration, to 5.6 cm and 2.7° after calibration. We emphasise
 453 that further calibration with a larger number of control parameters would enable further reduction in these errors.
 454 However, the OED framework tells us that increasing the number of control parameters is not justified, given the
 455 observation data used to constrain the unknown parameters.

456 6. Application of OED framework to a model of the northwest European continental shelf

457 In this section, we present the application of the OED framework to another model domain in order to further
 458 demonstrate the power of the method. We choose to use a model of the northwest European continental shelf. The
 459 model mesh is shown in figure 12a, which also indicates the locations of 42 tide gauges where observation data is
 460 available. We force the model at its ocean boundaries with the M2, S2, K1 and O1 tidal constituents. Each model
 461 run spans approximately 50 days, with the final 14.77 days used for harmonic analysis of model outputs. As described
 462 in section 3, the OED framework proceeds as follows:

- 463 1. Define the observation dataset. We use the M2 and S2 amplitude and phase observations at each of the 42 tide
 464 gauges indicated in figure 12b, resulting in a total of 168 data points.
- 465 2. Select a model setup with some suitable initial estimate for the Manning coefficient (we use a uniform value
 466 of $n = 0.025 \text{ s m}^{-1/3}$). Using the adjoint model, compute the gradient of each model output (corresponding
 467 to the 168 observations) with respect to the fully spatially varying bottom friction parameter. This requires a
 468 total of 168 adjoint model runs.
- 469 3. Propose a decomposition of the model domain into small regions, within which the bottom friction coefficient
 470 will always be uniform. Here we divide our model domain into 19 blocks, as indicated in figure 12b. The
 471 number and shape of these blocks is somewhat arbitrary, and these 19 blocks were drawn by hand. We provide

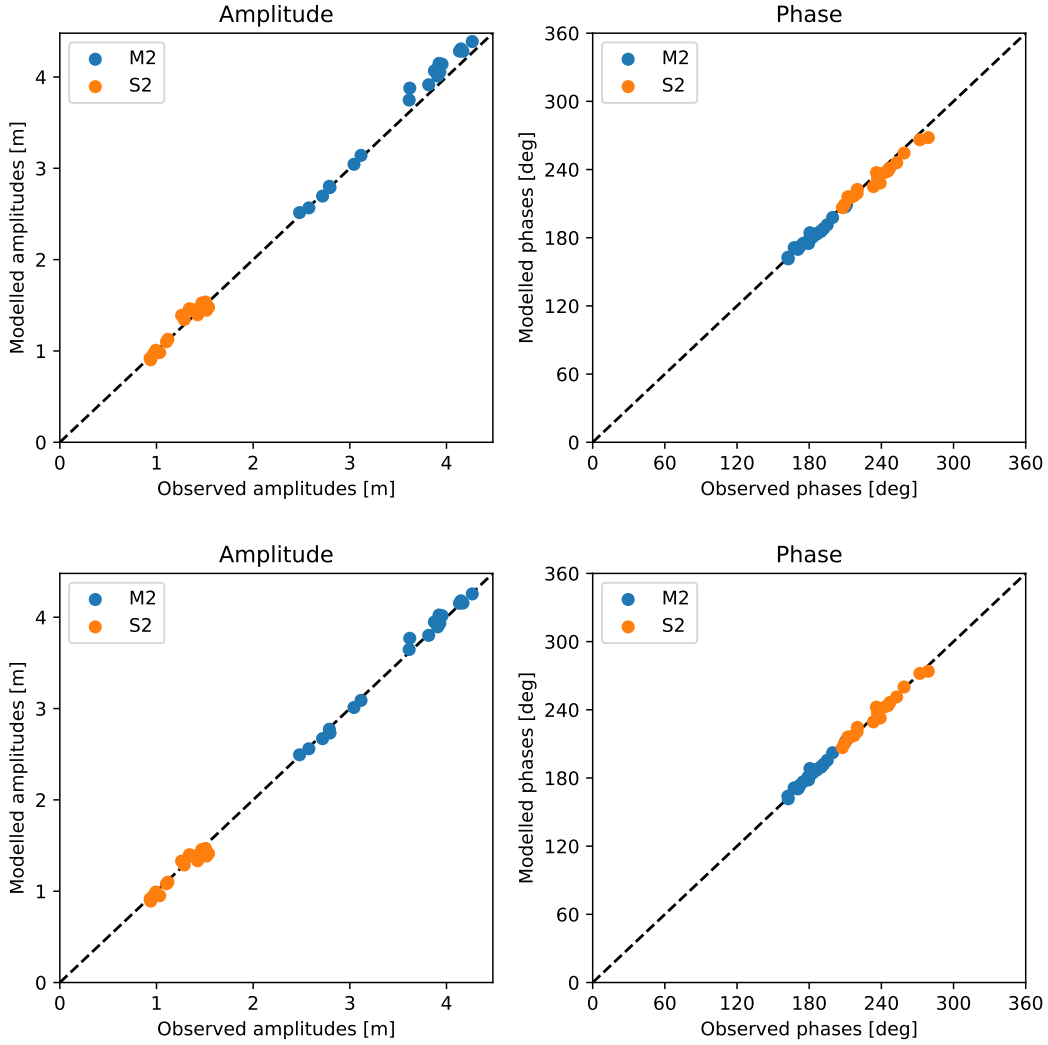


Figure 11: Scatter plots of modelled vs observed M2 and S2 amplitudes and phases in the Bristol Channel case study. Top: uncalibrated model with uniform friction coefficient $n = 0.025 \text{ s m}^{-1/3}$. Bottom: calibrated model, based on optimal $m = 3$ experiment design. The root mean squared errors are 9.7 cm and 4.1° before calibration, and 5.6 cm and 2.7° after calibration.

472 a further constraint that groupings of these subdomains must produce contiguous blocks. For a grouping into
 473 $m = 4$ parameters, this results in 31,865 possible experiment designs. Note that other possible approaches
 474 to parameterising the spatial variation of the bottom friction coefficient include the selection of independent
 475 points, between which the coefficient is determined by interpolation. The OED method can be trivially applied
 476 to any such space of experiment designs.

- 477 4. For each of the possible experiment designs, use the adjoint-derived gradients to compute the Fisher Information
 478 Matrix. Find the grouping which optimises a scalar measure of the FIM. Here we again use the modified D-
 479 criterion as described above.

480 The optimal experiment design, following the above schema for a grouping into $m = 4$ parameters, is shown in
 481 figure 13. The choice to use $m = 4$ is again somewhat arbitrary, and used here purely as a demonstration of the
 482 method. Note the large variation in the size of the parameter regions, and the number of tide gauges included within

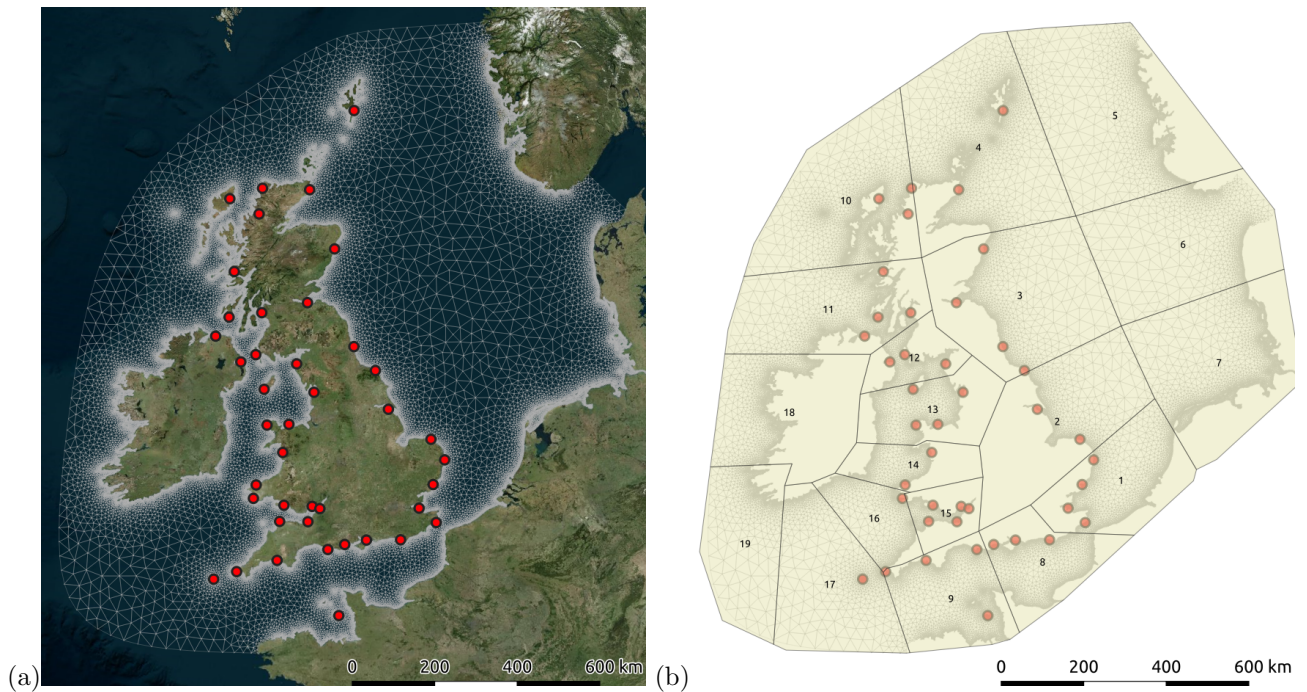


Figure 12: (a) Mesh used for continental shelf model. (b) Subdomains for friction parameterisation. The red circles indicate the locations of the tide gauges.

each. In particular, we find that the region including the North Sea covers by far the largest area, and contains a large number of tide gauges. Intuition might suggest that the optimal experiment design would likely consist of parameter regions of similar area, or containing similar volumes of observation data. However, a recent study on storm surge sensitivity to bottom friction coefficient in the North Sea (Warder et al., 2021) found that, due to the dynamics of surge propagation along the east coast of the UK, several tide gauge locations in the region exhibit similar spatial patterns of sensitivity to the bottom friction coefficient in the North Sea. This suggests that the tide gauges in the North Sea provide redundant information, which may explain why the optimal experiment design assigns a single friction parameter to the entire North Sea. In contrast, the dynamics within the Irish Sea are more complex, and the tide gauges in the region provide more diverse information about the bottom friction parameter, facilitating an experiment design featuring a greater number of friction parameters in the area.

7. Discussion

This study has demonstrated the application of an optimal experiment design procedure to the problem of bottom friction parameter estimation. In particular, we have addressed the problem that the spatial variation of the bottom friction parameter is unknown. The proposed OED method is able to identify a low-dimensional representation of the friction coefficient, resulting in parameters which can be constrained to a specified uncertainty by a given set of observations. We emphasise that the OED method is distinct from, and more useful than, the selection of a parameter space which can attain a specified agreement between model outputs and observations. A model calibration process motivated only by minimising the misfit between model outputs and observations would favour the use of a large number of tuning parameters, which can produce low model-observation misfits while uncertainty in the estimated parameters (and therefore in model outputs at new locations) remains high. A common technique

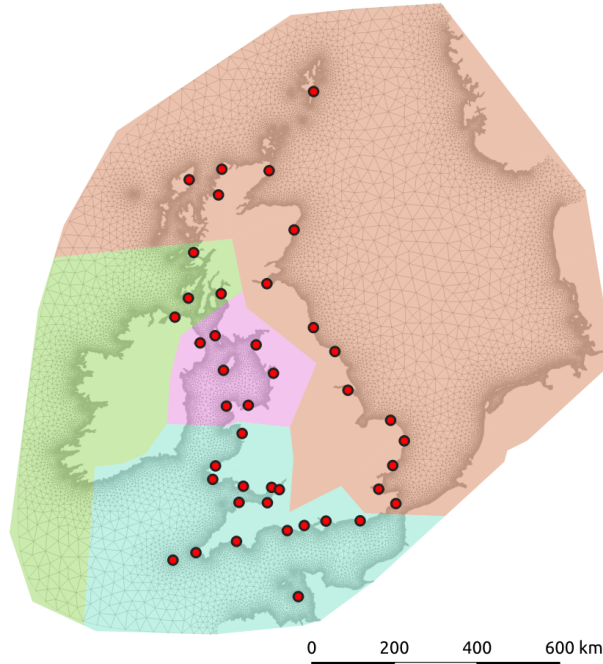


Figure 13: Optimal grouping of the subdomains into $m = 4$ groups.

503 within the literature for addressing this issue is regularisation of the parameter estimation problem. The OED
 504 framework proposed within this paper offers an alternative approach, avoiding the need for regularisation. The OED
 505 approach constitutes a rigorous method for identifying a parameterisation of the unknown friction coefficient field
 506 which can be constrained to a specified precision by a given set of observations.

507 However, we note that it is not possible to interpret the OED-derived spatial pattern of the friction coefficient
 508 field as providing information relating to the underlying physical processes governing bottom friction. Similarly, the
 509 poorer performance of experiment designs based on the *a priori* use of sediment data compared with simple slice-
 510 subdomains described in sections 4.1 and 4.2 does not imply that the observations are inconsistent with the sediment
 511 data. This is because, in selecting the bottom friction coefficient as the only model calibration parameter, we are
 512 implicitly assuming that bottom friction is the only source of model uncertainty. However, in reality there exist a
 513 variety of sources of modelling errors, including other uncertain model inputs, the choice of governing equations,
 514 and model discretisation errors. Since the bottom friction parameterisation accounts for subgrid-scale bathymetric
 515 variation, the mesh resolution is likely to be a particularly influential modelling choice. Despite this issue, it is
 516 common practice to perform model calibration with respect to bottom friction alone. The presented OED approach
 517 is therefore a valuable preliminary step in a model calibration study, since a tighter constraint on a selected set of
 518 unknown parameters increases confidence in model outputs at new locations, even if the calibrated parameters are
 519 implicitly correcting for other modelling errors.

520 Our results indicate that the number of friction parameters that can be well constrained by the observations is
 521 typically much smaller than the number of observations. This is due in part to the ratio between the observation error
 522 variance and the target parameter estimate covariance (and the relationship between them via the model), but also
 523 due to redundancy in the information contained in the observations. This is particularly evident in the application
 524 to the northwest European continental shelf. In other words, a large quantity of data does not justify the estimation

525 of a large number of model parameters, and the methods proposed within this work offer a rigorous technique for the
526 selection of an appropriate set of model parameters for estimation from a given set of observations. The results of
527 section 4.2 also demonstrate that the careful selection of the spatial distribution of the friction parameters (i.e. not
528 just the number of parameters) is important. We have also found that, for the model domain considered, the choice
529 to divide the domain by simple slices produces better-performing experiment designs than the use of sediment data
530 to *a priori* constrain the spatial variation of the bottom friction coefficient. However, as noted above, this can be
531 attributed to the relative sizes of the spaces of experiment designs for these choices, and does not necessarily imply
532 that additional physical knowledge via sediment data cannot be beneficial within parameter estimation.

533 It is useful to compare the computational cost associated with the OED method presented here with the cost of a
534 ‘trial and error’ approach where the parameter estimation problem is solved for a selection of candidate experiment
535 designs. The adjoint model runs used here for the construction of the Jacobian were a one-off computational
536 overhead. An adjoint run of the *Thetis* model used within this work takes around 2.4 times the wallclock time of
537 a forward model run. The 80 pairs of forward and adjoint runs required therefore incurred a computational cost of
538 around 270 forward model runs. In contrast, the Bayesian inference algorithm (via a Gaussian process emulator)
539 required 30, 40 and 50 model runs for two-, three- and four-parameter experiment designs, respectively. For the
540 problems considered here, therefore, the computational cost of the OED methodology is less than 10 times that of
541 an individual parameter estimation experiment. Since the number of candidate designs considered within this study
542 was of order 100, we therefore conclude that the adjoint-based OED method is computationally efficient compared
543 with the trial-and-error approach. We note that the cost of the adjoint-based OED approach scales linearly with
544 the number of observations used, and may therefore be poorly suited to an application to timeseries observations,
545 or an extension to a large number of tidal constituents. However, as shown in section 4.3, for the right choice of *a*
546 *priori* constraint on the experiment design space, the optimal design itself shows low sensitivity to the precise set of
547 observations used. An approach to adjoint-based OED for larger numbers of observations may therefore be to perform
548 the OED based on a sample of the total observations; the results of this work suggest that the resulting experiment
549 design may still be close to optimal, particularly if the observations contain significant redundant information.
550 Furthermore, the computational cost of the adjoint-based approach scales well with the space of experiment designs.
551 For applications where the number of observations is relatively small compared to the number of possible parameter
552 space configurations, the adjoint-based OED procedure presented here is a highly efficient approach. For example,
553 a relaxation of the relatively strong *a priori* constraints placed on the experiment designs within this work could
554 significantly increase the number of possible designs, without having a significant impact on the computational cost
555 of the study. The efficiency of the OED framework was also evident in the application to the northwest European
556 continental shelf model, where a space of over 30,000 experiment designs was explored.

557 We note that an additional aspect relating to the computational cost of the OED approach of this work is the
558 search algorithm used for finding the optimum within the space of possible experiment designs. Here, the selection of
559 *a priori* constraints resulting in a fairly modest space of possible designs ensured that an exhaustive search algorithm
560 was feasible. These *a priori* constraints also assisted in the interpretability of the optimal designs in terms of areas
561 of the model domain, which can be a useful feature of parameter selection methods (Kravaris et al., 2013). However,
562 the exploration of larger experiment design spaces will be considered in future work (in order to fully exploit the

563 advantages of the adjoint-based approach as described above) and may require alternative optimisation approaches,
564 such as genetic algorithms (e.g. Chu and Hahn (2007)), which have been used more commonly in the related problem
565 of OED with respect to the set of observations (e.g. Catania and Paladino (2009)).

566 The use of the OED framework can also be compared with more traditional approaches in the literature, where
567 low-dimensional parameterisations are selected less methodically. In such applications, regularisation (via a penalty
568 term added to the misfit functional) is typically required in order to avoid an under-constrained calibration problem,
569 where the observation data is insufficient to infer all of the unknown parameters. The use of the OED framework
570 presented within this paper avoids the under-constrained problem entirely, and can be considered as an alternative
571 to regularisation. The regularisation approach requires the selection of one or more regularisation parameters, con-
572 trolling the magnitude of the penalty term in the functional. A cross-validation method for regularisation parameter
573 selection, such as in Ullman and Wilson (1998), not only requires a second observation dataset for the cross-validation
574 (whereas the OED framework does not), but also requires the repeated optimisation of the model with respect to
575 the control parameters, for each possible choice of regularisation parameter, thus increasing the computational cost.
576 The construction and solution of the OED problem within our Bristol Channel case study required 80 adjoint model
577 runs (and note that these model runs can be performed simultaneously, i.e. the solution of the OED problem is
578 “embarrassingly parallel”). A typical adjoint gradient-based approach to model calibration might require $\mathcal{O}(20)$ iter-
579 ations to converge, and must be run several times for different regularisation parameter values. The OED approach
580 is therefore competitive in terms of computational cost.

581 There are a number of avenues for the further application of OED methods within coastal ocean model paramete-
582 ter estimation. Firstly, we chose within this work to construct the experiment designs based on subdomains, within
583 which the parameter is taken to be uniform. An alternative approach would be to use the OED framework to find
584 independent points, between which the parameter is determined by interpolation. Secondly, the choice within this
585 work to define a modified D-criterion to compare experiment designs has produced good results, but other design cri-
586 teria are possible. For example, Machado et al. (2009) combine a normalised D-criterion with a modified E-criterion
587 which characterises the shape of the parameter estimate confidence region, thus favouring designs producing similar
588 constraints on all unknown parameters (i.e. a more spherical parameter confidence ellipsoid). Thirdly, a complemen-
589 tary application of OED within friction parameter estimation would be to identify an optimal observation strategy
590 for constraining a given set of friction parameters. Integrated approaches to simultaneously optimise the observation
591 strategy and the parameter space are also possible (Chu and Hahn, 2008), and may be valuable in coastal ocean
592 applications. Fourthly, this study has considered only the estimation of uncertain bottom friction parameters. Al-
593 though this is a common parameter for calibration within the numerical coastal ocean modelling literature, parameter
594 estimation methods can also be applied to other model inputs including bathymetry (e.g. Moure et al. (2004)),
595 boundary conditions (e.g. Chen et al. (2014)), or combinations of multiple inputs simultaneously (e.g. Heemink et al.
596 (2002)). The OED framework presented here is highly general and can be readily applied to these other sources of
597 uncertainty. Note also that the computational cost of the OED methodology would not increase if additional model
598 inputs were included. Finally, while model nonlinearity does not appear to have prevented the identification of ‘good’
599 experiment designs within this work, the application of the OED method to other model inputs, or combinations
600 thereof, may require appropriate treatment of this nonlinearity, for example via maximin design (e.g. Ushijima and

601 Yeh (2015)). Several of these aspects will be explored in future work.

602 **8. Conclusion**

603 In this study, we have applied an optimal experiment design technique to the identification of optimal piecewise-
604 constant representations of the Manning's n coefficient within a numerical coastal ocean model. We have taken a
605 region containing the Bristol Channel and Severn Estuary as a primary case study, where we have used harmonic
606 analysis data at 20 locations within the domain for the estimation of the bottom friction parameter. We have further
607 demonstrated the application of the framework to a model of the northwest European continental shelf, based on
608 the assimilation of harmonic analysis data from 42 tide gauge locations. The key advantage of OED methods is that
609 the experiment design (i.e. the observation strategy or the configuration of the input parameter space for model
610 calibration) can be optimised in advance of the observations being made, or computational resources being applied
611 to perform the model calibration itself. Here we used a method based on the Fisher Information Matrix, with the
612 sensitivity of the model outputs with respect to the friction coefficient computed via a numerical adjoint model. As
613 the optimality criterion, we proposed a modified D-criterion related to the determinant of the Fisher Information
614 Matrix, motivated by achieving the tightest possible constraints on the estimated parameters, and avoiding the need
615 for regularisation of the calibration problem.

616 The results of this study suggest that the modified D-criterion performs well in characterising the ability of a
617 given experiment design to constrain the unknown friction parameters. This was verified by testing a variety of
618 experiment designs within a Bayesian inference algorithm, and comparing the ModD criterion with an equivalent
619 measure of the parameter covariance matrix returned by the Bayesian inference algorithm. The results show that
620 the experiment design has a strong influence on the parameter constraints achievable from a given set of observation
621 data. For example, the best experiment designs for the estimation of four parameters are able to provide tighter
622 constraints on the unknown parameters than poor designs estimating only three parameters. This demonstrates the
623 value of such an OED framework.

624 The framework also facilitates an investigation of the response of experiment design performance to the availability
625 of observation data. For the observation dataset used for the Bristol Channel case study, consisting of M2 and S2
626 harmonic amplitudes and phases at 20 gauge locations, we find that the observations contain a significant amount
627 of redundant information. The OED criterion initially increases rapidly as we increase the number of observation
628 locations used, but subsequently flattens. Once this occurs, the parameter constraints can only be improved by
629 a reduction in observation uncertainty, or the inclusion of additional types of observations, i.e. harmonic phases in
630 addition to amplitudes, or additional harmonic constituents. This result has implications on strategies for observation
631 surveys, and optimal experiment design with respect to observation strategy will be considered in future work.

632 Finally, while the computational cost of our approach depends on the volume of observation data, our results
633 demonstrate that for practical parameter estimation problems the OED procedure can be considered a computa-
634 tionally efficient approach. For our case studies, the overall computational cost of the OED framework was on the
635 order of 10 times that of performing an individual parameter estimation experiment with a given design. The OED
636 framework is therefore not prohibitively expensive, compared with either a trial-and-error approach to finding the
637 optimal experiment design, or a more traditional regularisation approach requiring repeated optimisations in order

638 to select a regularisation parameter.

639 In summary, we have demonstrated that our OED framework can be used to formulate well-constrained calibration
640 problems, and can be a valuable preliminary step in a parameter estimation study.

641 **Acknowledgements**

642 We acknowledge funding from EPSRC under grants EP/R029423/1, EP/R511547/1 as well as the EPSRC Centre
643 for Doctoral Training in Fluid Dynamics across Scales (EP/L016230/1). We additionally acknowledge the Research
644 Computing Service at Imperial College London for HPC resources, and thank Mariana Clare for helpful comments
645 on the manuscript. This study uses data from the National Tidal and Sea Level Facility, provided by the British
646 Oceanographic Data Centre and funded by the Environment Agency.

647 **References**

- 648 T. A. A. Adcock, S. Draper, and T. Nishino. Tidal power generation—a review of hydrodynamic modelling. *Proceedings*
649 *of the Institution of Mechanical Engineers, Part A: Journal of Power and Energy*, 229(7):755–771, 2015.
- 650 J. E. Alaña and C. Theodoropoulos. Optimal location of measurements for parameter estimation of distributed
651 parameter systems. *Computers & chemical engineering*, 35(1):106–120, 2011.
- 652 M. Altaf, M. Verlaan, and A. Heemink. Efficient identification of uncertain parameters in a large-scale tidal model of
653 the european continental shelf by proper orthogonal decomposition. *International Journal for Numerical Methods*
654 *in Fluids*, 68(4):422–450, 2012.
- 655 A. Angeloudis and R. A. Falconer. Sensitivity of tidal lagoon and barrage hydrodynamic impacts and energy outputs
656 to operational characteristics. *Renewable Energy*, 114:337–351, 2017.
- 657 G. J. Arcement and V. R. Schneider. Guide for selecting manning’s roughness coefficients for natural channels and
658 flood plains, 1989.
- 659 A. Avdis, A. S. Candy, J. Hill, S. C. Kramer, and M. D. Piggott. Efficient unstructured mesh generation for marine
660 renewable energy applications. *Renewable Energy*, 116:842–856, 2018.
- 661 E. Balsa-Canto, A. A. Alonso, and J. R. Banga. Computational procedures for optimal experimental design in
662 biological systems. *IET systems biology*, 2(4):163–172, 2008.
- 663 F. G. Blanchet, P. Legendre, and D. Borcard. Forward selection of explanatory variables. *Ecology*, 89(9):2623–2632,
664 2008.
- 665 F. Catania and O. Paladino. Optimal sampling for the estimation of dispersion parameters in soil columns using an
666 iterative genetic algorithm. *Environmental Modelling & Software*, 24(1):115–123, 2009.
- 667 H. Chen, A. Cao, J. Zhang, C. Miao, and X. Lv. Estimation of spatially varying open boundary conditions for a
668 numerical internal tidal model with adjoint method. *Mathematics and Computers in Simulation*, 97:14–38, 2014.
- 669 W.-B. Chen and W.-C. Liu. Investigating the fate and transport of fecal coliform contamination in a tidal estuarine
670 system using a three-dimensional model. *Marine Pollution Bulletin*, 116(1-2):365–384, 2017.
- 671 Y. Chu and J. Hahn. Parameter set selection for estimation of nonlinear dynamic systems. *AIChE journal*, 53(11):
672 2858–2870, 2007.
- 673 Y. Chu and J. Hahn. Integrating parameter selection with experimental design under uncertainty for nonlinear
674 dynamic systems. *AIChE Journal*, 54(9):2310–2320, 2008.
- 675 Y. Chu and J. Hahn. Parameter set selection via clustering of parameters into pairwise indistinguishable groups of
676 parameters. *Industrial & Engineering Chemistry Research*, 48(13):6000–6009, 2009.
- 677 S. Das and R. Lardner. On the estimation of parameters of hydraulic models by assimilation of periodic tidal data.
678 *Journal of Geophysical Research: Oceans*, 96(C8):15187–15196, 1991.

679 A. Davies and P. Robins. Residual flow, bedforms and sediment transport in a tidal channel modelled with variable
680 bed roughness. *Geomorphology*, 295:855–872, 2017.

681 A. de Brauwere, F. De Ridder, O. Gourgue, J. Lambrechts, R. Comblen, R. Pintelon, J. Passerat, P. Servais,
682 M. Elskens, W. Baeyens, et al. Design of a sampling strategy to optimally calibrate a reactive transport model:
683 Exploring the potential for escherichia coli in the scheldt estuary. *Environmental modelling & software*, 24(8):
684 969–981, 2009.

685 Digimap. Marine Themes Digital Elevation Model 6 Arc Second [ASC geospatial data], Scale 1:250000, Tiles:
686 2052010080, 2052010060, 2052010040, 2050010100, 2050010080, 2050010060, 2050010040, 2048010100, 2048010080,
687 2048010060, 2048010040, Updated: 25 October 2013, OceanWise, Using: EDINA Marine Digimap Service,
688 <https://digimap.edina.ac.uk>, Downloaded: 2020-07-14, 2013.

689 K. Döös, J. Nycander, and P. Sigray. Slope-dependent friction in a barotropic model. *Journal of Geophysical Research:*
690 *Oceans*, 109(C1), 2004.

691 G. D. Egbert and S. Y. Erofeeva. Efficient inverse modeling of barotropic ocean tides. *Journal of Atmospheric and*
692 *Oceanic Technology*, 19(2):183–204, 2002. doi: 10.1175/1520-0426(2002)019<0183:EIMOBO>2.0.CO;2.

693 A. F. Emery and A. V. Nenarokomov. Optimal experiment design. *Measurement Science and Technology*, 9(6):864,
694 1998.

695 P. E. Farrell, D. A. Ham, S. W. Funke, and M. E. Rognes. Automated derivation of the adjoint of high-level transient
696 finite element programs. *SIAM Journal on Scientific Computing*, 35(4):C369–C393, 2013.

697 V. V. Fedorov. *Theory of optimal experiments*. 1972.

698 R. A. Flather. Existing operational oceanography. *Coastal Engineering*, 41(1-3):13–40, 2000.

699 I. Ford, D. Titterton, and C. P. Kitsos. Recent advances in nonlinear experimental design. *Technometrics*, 31(1):
700 49–60x, 1989.

701 C. Geuzaine and J. F. Remacle. Gmsh: A 3-D finite element mesh generator with built-in pre- and post-processing
702 facilities. *International Journal for Numerical Methods in Engineering*, 79(11):1309–1331, 2009. doi: 10.1002/
703 nme.2579.

704 Z. Goss, D. Coles, and M. Piggott. Identifying economically viable tidal sites within the alderney race through
705 optimization of leveled cost of energy. *Philosophical Transactions of the Royal Society A*, 378(2178):20190500,
706 2020.

707 GPy. GPy: A Gaussian process framework in python. <http://github.com/SheffieldML/GPy>, since 2012.

708 L. Graham, T. Butler, S. Walsh, C. Dawson, and J. J. Westerink. A Measure-Theoretic Algorithm for Estimating
709 Bottom Friction in a Coastal Inlet: Case Study of Bay St. Louis during Hurricane Gustav (2008). *Monthly Weather*
710 *Review*, 145:929–954, 2017. ISSN 0027-0644. doi: 10.1175/mwr-d-16-0149.1.

- 711 N. Guillou and J. Thiébot. The impact of seabed rock roughness on tidal stream power extraction. *Energy*, 112:
712 762–773, 2016. doi: 10.1016/j.energy.2016.06.053.
- 713 J. W. Hall, L. J. Manning, and R. K. Hankin. Bayesian calibration of a flood inundation model using spatial data.
714 *Water Resources Research*, 47(5):1–14, 2011. doi: 10.1029/2009WR008541.
- 715 W. K. Hastings. Monte Carlo sampling methods using Markov chains and their applications. 1970.
- 716 A. Heemink, E. Mouthaan, M. Roest, E. Vollebregt, K. Robaczewska, and M. Verlaan. Inverse 3d shallow water flow
717 modelling of the continental shelf. *Continental Shelf Research*, 22(3):465–484, 2002.
- 718 P. Hristov, F. DiazDelaO, E. S. Flores, C. Guzmán, and U. Farooq. Probabilistic sensitivity analysis to understand
719 the influence of micromechanical properties of wood on its macroscopic response. *Composite Structures*, 181:
720 229–239, 2017.
- 721 X. Huan and Y. M. Marzouk. Simulation-based optimal bayesian experimental design for nonlinear systems. *Journal*
722 *of Computational Physics*, 232(1):288–317, 2013.
- 723 N. Huybrechts, H. Smaoui, S. Orseau, P. Tassi, and F. Klein. Automatic calibration of bed friction coefficients to
724 reduce the influence of seasonal variation: Case of the gironde estuary. *Journal of Waterway, Port, Coastal, and*
725 *Ocean Engineering*, 147(3):05021004, 2021.
- 726 T. Kärnä, B. De Brye, O. Gourgue, J. Lambrechts, R. Comblen, V. Legat, and E. Deleersnijder. A fully implicit
727 wetting–drying method for dg-fem shallow water models, with an application to the scheldt estuary. *Computer*
728 *Methods in Applied Mechanics and Engineering*, 200(5-8):509–524, 2011.
- 729 T. Kärnä, S. C. Kramer, L. Mitchell, D. A. Ham, M. D. Piggott, and A. M. Baptista. Thetis coastal ocean
730 model: discontinuous galerkin discretization for the three-dimensional hydrostatic equations. *Geoscientific Model*
731 *Development*, 11(11):4359–4382, 2018.
- 732 S. Kramer, T. Kärnä, J. Hill, and S. W. Funke. stephankramer/uptide: First release of uptide v1.0, 2020. [http:
733 //doi.org/10.5281/zenodo.3909652](http://doi.org/10.5281/zenodo.3909652).
- 734 C. Kravaris, J. Hahn, and Y. Chu. Advances and selected recent developments in state and parameter estimation.
735 *Computers & chemical engineering*, 51:111–123, 2013.
- 736 R. H. Langland. Issues in targeted observing. *Quarterly Journal of the Royal Meteorological Society: A journal of*
737 *the atmospheric sciences, applied meteorology and physical oceanography*, 131(613):3409–3425, 2005.
- 738 R. Lardner, A. Al-Rabeh, and N. Gunay. Optimal estimation of parameters for a two-dimensional hydrodynamical
739 model of the arabian gulf. *Journal of Geophysical Research: Oceans*, 98(C10):18229–18242, 1993.
- 740 R. Li, M. A. Henson, and M. J. Kurtz. Selection of model parameters for off-line parameter estimation. *IEEE*
741 *Transactions on control systems technology*, 12(3):402–412, 2004.
- 742 X. Li, A. Plater, and N. Leonardi. Modelling the transport and export of sediments in macrotidal estuaries with
743 eroding salt marsh. *Estuaries and coasts*, 41(6):1551–1564, 2018.

- 744 X. Lu and J. Zhang. Numerical study on spatially varying bottom friction coefficient of a 2d tidal model with adjoint
745 method. *Continental Shelf Research*, 26(16):1905–1923, 2006.
- 746 C. Lyddon, J. M. Brown, N. Leonardi, and A. J. Plater. Uncertainty in estuarine extreme water level predictions
747 due to surge-tide interaction. *PloS one*, 13(10):e0206200, 2018.
- 748 V. C. Machado, G. Tapia, D. Gabriel, J. Lafuente, and J. A. Baeza. Systematic identifiability study based on
749 the fisher information matrix for reducing the number of parameters calibration of an activated sludge model.
750 *Environmental Modelling & Software*, 24(11):1274–1284, 2009.
- 751 L. Mackie, D. Coles, M. Piggott, and A. Angeloudis. The Potential for Tidal Range Energy Systems to Provide
752 Continuous Power: A UK Case Study. *Journal of Marine Science and Engineering*, 8(10):780, 2020a.
- 753 L. Mackie, P. S. Evans, M. J. Harrold, T. O’Doherty, M. D. Piggott, and A. Angeloudis. Modelling hydrodynamics
754 in an energetic tidal strait with pronounced bathymetric features. *In review, submitted to Applied Ocean Research*,
755 2020b. doi: <https://doi.org/10.31223/osf.io/8txmd>.
- 756 K. N. Marshall, I. C. Kaplan, E. E. Hodgson, A. Hermann, D. S. Busch, P. McElhany, T. E. Essington, C. J. Harvey,
757 and E. A. Fulton. Risks of ocean acidification in the california current food web and fisheries: ecosystem model
758 projections. *Global Change Biology*, 23(4):1525–1539, 2017.
- 759 S. Maßmann. *Tides on unstructured meshes*. PhD thesis, Universitat Bremen, 2010.
- 760 T. Mayo, T. Butler, C. Dawson, and I. Hoteit. Data assimilation within the Advanced Circulation (ADCIRC)
761 modeling framework for the estimation of Manning’s friction coefficient. *Ocean Modelling*, 76:43–58, 2014. ISSN
762 14635003. doi: 10.1016/j.ocemod.2014.01.001. URL <http://dx.doi.org/10.1016/j.ocemod.2014.01.001>.
- 763 S. K. Mitusch, S. W. Funke, and J. S. Dokken. dolfin-adjoint 2018.1: automated adjoints for FEniCS and Firedrake.
764 *Journal of Open Source Software*, 4(38), 2019.
- 765 B. Moure, P. De Mey, F. Lyard, and C. Le Provost. Assimilation of sea level data over continental shelves: an
766 ensemble method for the exploration of model errors due to uncertainties in bathymetry. *Dynamics of atmospheres
767 and oceans*, 38(2):93–121, 2004.
- 768 S. P. Neill, A. Angeloudis, P. E. Robins, I. Walkington, S. L. Ward, I. Masters, M. J. Lewis, M. Piano, A. Avdis, M. D.
769 Piggott, et al. Tidal range energy resource and optimization—past perspectives and future challenges. *Renewable
770 energy*, 127:763–778, 2018.
- 771 W. Pan, S. C. Kramer, T. Kärnä, and M. D. Piggott. Comparing non-hydrostatic extensions to a discontinuous finite
772 element coastal ocean model. *Ocean Modelling*, 151:101634, 2020. ISSN 1463-5003. doi: [https://doi.org/10.1016/
773 j.ocemod.2020.101634](https://doi.org/10.1016/j.ocemod.2020.101634). URL <http://www.sciencedirect.com/science/article/pii/S1463500319303221>.
- 774 A. Pepelyshev, V. B. Melas, N. Strigul, and H. Dette. Design of experiments for the monod model: robust and
775 efficient designs. Technical report, Technical Report, 2004.

- 776 R. Periañez, M. Casas-Ruíz, and J. Bolívar. Tidal circulation, sediment and pollutant transport in cádis bay (sw
777 spain): a modelling study. *Ocean engineering*, 69:60–69, 2013.
- 778 C. E. Rasmussen. Gaussian processes in machine learning. In *Summer School on Machine Learning*, pages 63–71.
779 Springer, 2003.
- 780 F. Rathgeber, D. A. Ham, L. Mitchell, M. Lange, F. Luporini, A. T. McRae, G.-T. Bercea, G. R. Markall, and
781 P. H. Kelly. Firedrake: automating the finite element method by composing abstractions. *ACM Transactions on*
782 *Mathematical Software (TOMS)*, 43(3):1–27, 2016.
- 783 C. R. Rojas, J. S. Welsh, G. C. Goodwin, and A. Feuer. Robust optimal experiment design for system identification.
784 *Automatica*, 43(6):993–1008, 2007.
- 785 SHOM, 2013. URL <https://data.shom.fr>. Accessed 2019-04-10.
- 786 I. M. Sobol. Global sensitivity indices for nonlinear mathematical models and their Monte Carlo estimates. *Mathe-*
787 *matics and Computers in Simulation*, 55(1-3):271–280, 2001.
- 788 T. Söderström and P. Stoica. *System identification*. Prentice-Hall International, 1989.
- 789 I. Sraj, M. Iskandarani, A. Srinivasan, W. C. Thacker, J. Winokur, A. Alexanderian, C. Y. Lee, S. S. Chen, and
790 O. M. Knio. Bayesian inference of drag parameters using AXBT data from typhoon fanapi. *Monthly Weather*
791 *Review*, 141(7):2347–2367, 2013. doi: 10.1175/MWR-D-12-00228.1.
- 792 I. Sraj, M. Iskandarani, W. C. Thacker, A. Srinivasan, and O. M. Knio. Drag parameter estimation using gradients
793 and hessian from a polynomial chaos model surrogate. *Monthly Weather Review*, 142(2):933–941, 2014a.
- 794 I. Sraj, K. T. Mandli, O. M. Knio, C. N. Dawson, and I. Hoteit. Uncertainty quantification and inference of Manning’s
795 friction coefficients using DART buoy data during the Tōhoku tsunami. *Ocean Modelling*, 83:82–97, 2014b. doi:
796 10.1016/j.ocemod.2014.09.001.
- 797 N. Strigul, H. Dette, and V. B. Melas. A practical guide for optimal designs of experiments in the monod model.
798 *Environmental Modelling & Software*, 24(9):1019–1026, 2009.
- 799 A. Y. Sun. A robust geostatistical approach to contaminant source identification. *Water Resources Research*, 43(2),
800 2007.
- 801 D. Ucinski. *Optimal measurement methods for distributed parameter system identification*. CRC press, 2004.
- 802 D. S. Ullman and R. E. Wilson. Model parameter estimation from data assimilation modeling: Temporal and spatial
803 variability of the bottom drag coefficient. *Journal of Geophysical Research: Oceans*, 103(C3):5531–5549, 1998.
- 804 T. T. Ushijima and W. W. Yeh. Experimental design for estimating unknown hydraulic conductivity in an aquifer
805 using a genetic algorithm and reduced order model. *Advances in Water Resources*, 86:193–208, 2015.
- 806 V. Vandenberghe, A. van Griensven, and W. Bauwens. Detection of the most optimal measuring points for water
807 quality variables: application to the river water quality model of the river dender in eswat. *Water science and*
808 *technology*, 46(3):1–7, 2002.

- 809 C. V. Vouriot, A. Angeloudis, S. C. Kramer, and M. D. Piggott. Fate of large-scale vortices in idealized tidal lagoons.
810 *Environmental Fluid Mechanics*, 19(2):329–348, 2019.
- 811 É. Walter and L. Pronzato. Qualitative and quantitative experiment design for phenomenological models—a survey.
812 *Automatica*, 26(2):195–213, 1990.
- 813 D. Wang, A. Cao, J. Zhang, D. Fan, Y. Liu, and Y. Zhang. A three-dimensional cohesive sediment transport model
814 with data assimilation: Model development, sensitivity analysis and parameter estimation. *Estuarine, Coastal and*
815 *Shelf Science*, 206:87–100, 2018.
- 816 T. Wang and Z. Yang. A tidal hydrodynamic model for cook inlet, alaska, to support tidal energy resource charac-
817 terization. *Journal of Marine Science and Engineering*, 8(4):254, 2020.
- 818 S. C. Warder, K. J. Horsburgh, and M. D. Piggott. Understanding the contribution of uncertain wind stress to storm
819 surge predictions. In *4th IMA International Conference on Flood Risk, Swansea*, 2019.
- 820 S. C. Warder, K. J. Horsburgh, and M. D. Piggott. Adjoint-based sensitivity analysis for a numerical storm surge
821 model. *Ocean Modelling*, 160:101766, 2021.
- 822 P. Wessel and W. H. F. Smith. A global, self-consistent, hierarchical, high-resolution shoreline database. *Journal of*
823 *Geophysical Research: Solid Earth*, 101(B4):8741–8743, 1996. doi: 10.1029/96JB00104.
- 824 P. Whomersley, J. Van der Molen, D. Holt, C. Trundle, S. Clark, and D. Fletcher. Modeling the dispersal of
825 spiny lobster (*palinurus elephas*) larvae: Implications for future fisheries management and conservation measures.
826 *Frontiers in Marine Science*, 5:58, 2018.
- 827 H. Yu, H. Yue, and P. Halling. Comprehensive experimental design for chemical engineering processes: A two-layer
828 iterative design approach. *Chemical Engineering Science*, 189:135–153, 2018.
- 829 J. Zhang, X. Lu, P. Wang, and Y. P. Wang. Study on linear and nonlinear bottom friction parameterizations for
830 regional tidal models using data assimilation. *Continental Shelf Research*, 31(6):555–573, 2011.

831 **Appendix A. Bayesian inference**

832 Bayesian inference is a powerful statistical technique for solving inverse problems, and has been applied to bottom
 833 friction parameter estimation previously (Hall et al., 2011, Sraj et al., 2014b). Within this work, we use Bayesian
 834 inference to solve parameter estimation problems for a selected number of experiment designs. Our Bayesian inference
 835 framework proceeds via a Markov Chain Monte Carlo (MCMC) method, using a Gaussian process emulator (GPE)
 836 as a surrogate for the full numerical model, since the method relies on a large number of model evaluations. Appendix
 837 A.1 describes Gaussian process emulation, and Appendix A.2 details the MCMC algorithm.

838 *Appendix A.1. Gaussian process emulation*

839 This exposition follows Rasmussen (2003), to which the reader is referred for further detail. The crux of Gaussian
 840 process emulation is that, under the assumption that model outputs follow a multivariate Gaussian distribution, a
 841 vector of ‘test’ model outputs \mathbf{f}_* for model inputs \mathbf{X}_* can be predicted from ‘training’ model outputs \mathbf{f} computed for
 842 model inputs \mathbf{X} . The model outputs satisfy the conditional distribution

$$\mathbf{f}_* | \mathbf{f} \sim \mathcal{N}(\mu(\mathbf{X}_*) + \Sigma_*^T \Sigma^{-1}(\mathbf{f} - \mu(\mathbf{X})), \Sigma_{**} - \Sigma_*^T \Sigma^{-1} \Sigma_*), \quad (\text{A.1})$$

843 where μ denotes the mean function, Σ the covariance matrix for the training points, Σ_{**} the covariance matrix for
 844 the test points, and Σ_* the covariance between the training and test points. The covariance matrices are typically
 845 parameterised by

$$\Sigma = k(\mathbf{X}, \mathbf{X}); \Sigma_* = k(\mathbf{X}, \mathbf{X}_*); \Sigma_{**} = k(\mathbf{X}_*, \mathbf{X}_*), \quad (\text{A.2})$$

846 where $k(x, x')$ is a covariance function, with a common choice given by

$$k(x, x') = \sigma^2 C(x, x'), \quad (\text{A.3})$$

847 where σ^2 is a covariance parameter and $C(x, x')$ a parameterised correlation function. Similarly, the mean function
 848 μ is typically taken as a simple linear or quadratic function of the model inputs. The parameters introduced by the
 849 mean and covariance functions can be determined via maximum likelihood estimation using the training data. Once
 850 these parameters are determined, equation (A.1) can be evaluated at low computational cost, and the mean of the
 851 conditional distribution estimates model outputs for unseen values of the model inputs. The generalisation to higher
 852 dimensions and multiple model outputs is straightforward. Within this work, we use the Python package GPy (since
 853 2012); the reader is referred to the package documentation for further implementation detail.

854 For the application to Bayesian inference within this study, the model inputs are the vector of unknown parameters
 855 $\boldsymbol{\theta}$ for a given experiment design, and the model outputs are the M2 and S2 amplitudes and phases at each of the 20
 856 tide gauge locations (a total of 80 outputs). The training input samples (\mathbf{X} in the above notation) are selected using
 857 Latin Hypercube Sampling from uniform prior distributions in the range [0.01, 0.05] for each Manning coefficient.
 858 According to the ‘10d’ rule (Sobol, 2001, Hristov et al., 2017), it is common to train a Gaussian process emulator
 859 using at least 10d samples, where d is the number of input parameters. As a cautious approach, within this work
 860 we use 10(d + 1) training samples for each selected experiment design to be tested using Bayesian inference. The
 861 corresponding output for each training sample is generated from a forward run of the Thetis numerical model
 862 comprising a 13-day spin-up period followed by a 14.77-day harmonic analysis period. Harmonic analysis for the M2

and S2 constituents is performed at each of the 20 tide gauge locations, using the Python package uptide (Kramer et al., 2020). These outputs are used as the training data to construct a Gaussian process emulator as described above.

Prior to the use of the GPEs within the Bayesian inference algorithm, we first verified their faithfulness as a surrogate for the full numerical model by comparing their outputs with additional full numerical model runs for random samples of the model parameters. We found that the GPE error covariance calculated via equation (A.1) provides a good estimate of the true emulator error covariance, and that this covariance is small compared with the assumed observation uncertainty. This covariance is therefore neglected when using the GPE within the Bayesian inference algorithm.

Appendix A.2. Markov Chain Monte Carlo algorithm

The Bayesian inference framework follows a similar approach to Sraj et al. (2013, 2014b). Denoting the vector of tidal harmonic observations \mathbf{y} and the vector of model parameters $\boldsymbol{\theta}$, Bayes' theorem gives

$$\Pi(\boldsymbol{\theta}|\mathbf{y}) \propto L(\mathbf{y}|\boldsymbol{\theta})q(\boldsymbol{\theta}), \quad (\text{A.4})$$

where Π is the posterior distribution of the parameters $\boldsymbol{\theta}$ given the observed data \mathbf{y} , L is the likelihood of observing the outputs \mathbf{y} given the parameters $\boldsymbol{\theta}$, and q is the prior distribution of the parameter vector $\boldsymbol{\theta}$. The prior distribution on each individual Manning coefficient θ_i is taken as uniform in the range $[0.01, 0.05]$, hence

$$q(\boldsymbol{\theta}) = \begin{cases} \left(\frac{1}{0.05-0.01}\right)^m & \text{if } 0.01 < \theta_i < 0.05 \ \forall i \\ 0 & \text{otherwise.} \end{cases} \quad (\text{A.5})$$

It is assumed that the model-observation discrepancies $\mathbf{y} - \mathbf{f}$, where \mathbf{f} is the vector of model outputs, are independent and identically distributed variables with zero mean, and a covariance matrix

$$\Sigma_{obs} = \begin{pmatrix} \sigma_{\text{amp}}^2 \mathbb{I}_{40 \times 40} & 0 \\ 0 & \sigma_{\text{phase}}^2 \mathbb{I}_{40 \times 40} \end{pmatrix}, \quad (\text{A.6})$$

where σ_{amp}^2 and σ_{phase}^2 are the amplitude and phase measurement covariances, respectively, as described in section 2.2, and \mathbb{I} denotes an identity matrix of the specified size (we use 40 amplitude observations and 40 phase observations). The likelihood $L(\mathbf{y}|\boldsymbol{\theta})$ is therefore given by

$$L(\mathbf{y}|\boldsymbol{\theta}) = (2\pi)^{-K/2} |\Sigma_{obs}|^{-1/2} \exp\left(-\frac{1}{2}(\mathbf{y} - \mathbf{f})^T \Sigma_{obs}^{-1} (\mathbf{y} - \mathbf{f})\right), \quad (\text{A.7})$$

where $K = 80$ is the number of observations.

Equation (A.4) gives the probability distribution of the unknown Manning coefficients, given the set of observations \mathbf{y} , and its evaluation represents the parameter estimation problem. A technique for sampling this posterior distribution when it cannot be directly calculated is the Markov Chain Monte Carlo (MCMC) method, which has the advantage that the constant of proportionality need not be determined. This work uses an implementation of the Random Walk Metropolis Hastings MCMC algorithm (Hastings, 1970), which is given by algorithm 1. Within the algorithm, we set the proposal distribution covariance matrix to

$$\Sigma_{\text{step}} = 0.001^2 \mathbb{I}_{m \times m}. \quad (\text{A.8})$$

890 This was found to give satisfactory results, without the need for an adaptive MCMC algorithm.

891 In the results presented within this paper, the number of MCMC samples is selected as $M = 10^6$, with the first
892 10^5 samples discarded as a burn-in period. The resulting chain of values $\boldsymbol{\theta}^{[i]}$ generated by the MCMC algorithm
893 constitute samples from the posterior distribution, from which the covariance of the estimated parameters can be
894 trivially computed.

Algorithm 1: Random Walk Metropolis Hastings algorithm

Initial guess for parameters $\boldsymbol{\theta} = \boldsymbol{\theta}^{[0]}$;

for $i = 1 : M$ **do**

1. Draw proposed set of parameters $\boldsymbol{\theta}^*$ from multivariate normal proposal distribution:

$$\boldsymbol{\theta}^* \sim \mathcal{N}(\boldsymbol{\theta}^{[i-1]}, \Sigma_{\text{step}})$$

2. Compute posterior $\Pi(\boldsymbol{\theta}^* | \mathbf{y})$ from Eq. (A.4)

3. Calculate $p_{\text{accept}} = \min\left(1, \frac{\Pi(\boldsymbol{\theta}^* | \mathbf{y})}{\Pi(\boldsymbol{\theta}^{[i-1]} | \mathbf{y})}\right)$

4. Generate $u \sim U(0, 1)$ and set $\boldsymbol{\theta}^{[i]} = \boldsymbol{\theta}^*$ if $p_{\text{accept}} > u$. Otherwise, set $\boldsymbol{\theta}^{[i]} = \boldsymbol{\theta}^{[i-1]}$.

end
