

Impact of injection rate ramp-up on nucleation and arrest of dynamic fault slip

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Abstract

Fluid injection into underground formations reactivates preexisting geological discontinuities such as faults or fractures. In this work, we investigate the impact of injection rate ramp-up present in many standard injection protocols on the nucleation and potential arrest of dynamic slip along a planar pressurized fault. We assume a linear increasing function of injection rate with time, up to a given time t_c after which a maximum value Q_m is achieved. Under the assumption of negligible shear-induced dilatancy and impermeable host medium, we solve numerically the coupled hydro-mechanical model and explore the different slip regimes identified via scaling analysis. We show that in the limit when fluid diffusion time scale t_w is much larger than the ramp-up time scale t_c , slip on an ultimately stable fault is essentially driven by pressurization at constant rate. Vice versa, in the limit when $t_c/t_w \gg 1$, the pressurization rate, quantified by the dimensionless ratio $\frac{Q_m t_w}{t_c Q^*}$, does impact both nucleation time and arrest distance of dynamic slip. Indeed, for a given initial fault loading condition and frictional weakening property, lower pressurization rates delay the nucleation of a finite-sized dynamic event and increase the corresponding run-out distance approximately proportional to $\propto \left(\frac{Q_m t_w}{t_c Q^*}\right)^{-0.472}$. On critically stressed faults, instead, the ramp-up of injection rate activates quasi-static slip which quickly turn into a run-away dynamic rupture. Its nucleation time decreases non-linearly with increasing value of $\frac{Q_m t_w}{t_c Q^*}$ and it may precede (or not) the one associated with fault pressurization at constant rate only.

Keywords: Fault slip, Nucleation, Dynamic rupture, Induced seismicity

1. Introduction

Anthropogenic fluid injection into underground formations is a common operation in many industrial applications. In the context of deep geothermal energy extraction, for instance, fluid is injected into targeted deep fault/fracture zones in order to enhance reservoir permeability and hence fluid circulation between injection and production wells (Giardini, 2009; Deichmann and Giardini, 2009). Among other applications that involve injection of fluid into subsurface there are hydraulic fracturing for oil and gas extraction from hydrocarbon reservoirs and wastewater disposal using deep wells (Warpinski and Teufel, 1987; Horton, 2012).

Although these engineering techniques are widely used, the injection of fluids into underground formations alters the local equilibrium of the Earth's crust, inducing micro-seismicity and, in some cases, large earthquakes (Horton, 2012; Ellsworth, 2013; Kim, 2013; Keranen et al., 2014; Weingarten et al., 2015). Significant earthquakes have been directly correlated with the injection activities, among these the Pohang earthquake of 2017 in South Korea with $M_w = 5.5$ (Kim et al., 2018; Grigoli et al., 2018; Yeo et al., 2020), the four $M_w = 3$ events in Basel, Switzerland, between 2006 and 2007 (Deichmann and Giardini, 2009; Goertz-Allmann et al., 2011), an event of $M_w = 3.5$ in the city of St. Gallen, Switzerland, back in 2013 (Diehl et al., 2014; Edwards et al., 2015; Diehl et al., 2017) and the $M_w = 5.7$ earthquake near Prague, Oklahoma, in 2011 (Keranen et al., 2013; Sumy et al., 2014).

Many numerical and theoretical models have been developed in order to investigate the impact of operational design parameters, such as injection pressure or injection rate, on fault slip activation and earthquakes nucleation upon fluid injection. Most of them are based on a Rate- and State-dependent friction model and, therefore, are well suited to explain features of earthquake cycles and seismicity rates. For example, Dempsey and Riffault (2019) used a pressure diffusion model coupled to R&S friction model to show that a reduction in injection rate may lead to a decrease in the seismicity rate in Oklahoma (USA). A similar result has been obtained by Lagenbruch and Zoback (2016) using instead a statistical model calibrated over many injection induced-earthquakes in Oklahoma. Using a poroelastic model incorporating R&S friction, Barbour et al. (2017) observed that in Oklahoma a variable injection rate may lead to a larger seismicity rate in-

36 creases compared to the one under constant injection rate (for an equivalent
37 injected volume). Chang et al. (2018), instead, studied the effect of injection
38 rate variation on seismicity rate post shut-in and they showed that a gradual
39 reduction of injection rate minimises the post-injection seismicity rate. Using
40 a Dietrich-Ruina heterogeneous 2D fault, Almakari et al. (2019) investigated
41 the effect of injection scenario not only in terms of seismicity rate, but also
42 in terms of magnitude content. They showed that the total seismic moment
43 increases with both maximum pressure and pressure rate and that the total
44 number of induced seismic events is controlled by the maximum pressure.
45 A recent study of Rudnicki and Zhan (2020) on a spring-block model shows
46 that larger pressurization rates stabilize fault slip events due to rate and state
47 friction.

48 The role of injection design parameters on fault slip behaviour has been ex-
49 tensively investigated also in many laboratory experiments. Among others,
50 Wang et al. (2020) showed that fault slip propagation is mainly governed by
51 fluid pressurization rate rather than injection pressure. French et al. (2016),
52 instead, observed that fluid pressurization is less effective than mechanical
53 changes in the fault normal stress at initiating accelerated slip events. The
54 effect of fluid pressure oscillations on fault slip stability has been investigated
55 by Noël et al. (2019) via a triaxial laboratory experiment. They showed that
56 perturbations caused by pore fluid oscillations promote seismic slip and that
57 seismic activity along a fault increases for increasing oscillation's amplitudes.
58 Despite the numerous studies on the effects of injection parameters on earth-
59 quakes nucleation and occurrence along faults, the impact of pressurization
60 rate on the onset of dynamic fault slip remains still elusive. Garagash and
61 Germanovich (2012) investigated extensively the conditions of nucleation and
62 arrest of dynamic fault slip on a frictional weakening pressurized fault. Their
63 generic findings, however, are valid only for two types of fault pressuriza-
64 tions, constant over-pressure and constant injection rate, which are over-
65 simplification of many injection protocols commonly used in industrial appli-
66 cations. In numerous fault reactivation experiments, for instance related to
67 deep geothermal energy exploitation, the injection protocol consists of one
68 (or more) stimulation cycle, where the controlled injection rate or injection
69 pressure increases in time (typically with stair-like increments), up to reach
70 a steady state regime, followed then by a shut-in phase (Hofmann et al.,
71 2018). In the hydraulic stimulation experiments conducted in Grimsel Test
72 Site, Switzerland, back in 2017 for example, the injection protocol consisted
73 of 4 injection cycles in which either injection pressure or injection flow rate

74 was increased in a stepwise manner, before reaching a plateau and shut-in
75 phase (Villiger et al., 2020). A similar trend of injection rate was also used
76 during the hydraulic stimulation of Enhanced Geothermal System in the city
77 of Basel, Switzerland (2006) (Häring et al., 2008) and in the more recent
78 EGS project in Pohang, South Korea (Hofmann et al., 2018), to cite a few
79 examples.

80 In this contribution, we extend the model of Garagash and Germanovich
81 (2012) to account for an initial ramp-up of injection rate in time and investi-
82 gate its effect on dynamic fault slip nucleation and arrest. We approximate
83 the step-wise increase of injection rate adopted in standard injection pro-
84 tocols using a simple linear increasing function with time, followed by a
85 maximum plateau (the shut-in phase is out of context and thus is not con-
86 sidered in this work). This choice represents a good approximation when the
87 time scale of each increment is much larger than the corresponding one of
88 each step (and thus a linear increasing function is a reasonable approxima-
89 tion). We solve the two-dimensional hydro-mechanical problem numerically
90 and verify the results with theoretical predictions. The goal of this study is
91 to evaluate the effects of the pressurization rate, quantified by the injection
92 rate ramp-up variation, on the nucleation and arrest of a seismic rupture on
93 a frictional weakening planar fault.

94 2. Fault model

95 We consider a planar fault embedded in an isotropic, homogeneous and
96 unbounded elastic medium under plane-strain conditions (see Figure 1). The
97 fault is subjected to an ambient pore-pressure p_o and a uniform far-field
98 stress state that resolved on the fault plane result in an effective normal
99 $\sigma'_o = \sigma_n - p_o$ and shear τ^o stress component. Such a uniform ambient stress
100 state, typical of a limited fault extent compared to the background in-situ
101 gradient, is perturbed via a point source injection of volumetric flow rate
102 $Q(t)$ [L^2/T] directly in the middle of the fault (specifically at $x = 0$). In
103 order to investigate the effect of injection rate ramp-up on fault slip stability,
104 we consider a linear increase of injection flow rate in time, followed by a
105 plateau after a given time t_c (the shut-in phase is out of scope here, hence
106 it will not be considered - see Figure 1). This parametrisation represents
107 an approximation of many standard fault injection protocols used in hydro-
108 shearing stimulation of fractured reservoirs, in which the design (constant)
109 value of injection rate is reached upon stair-like increments.

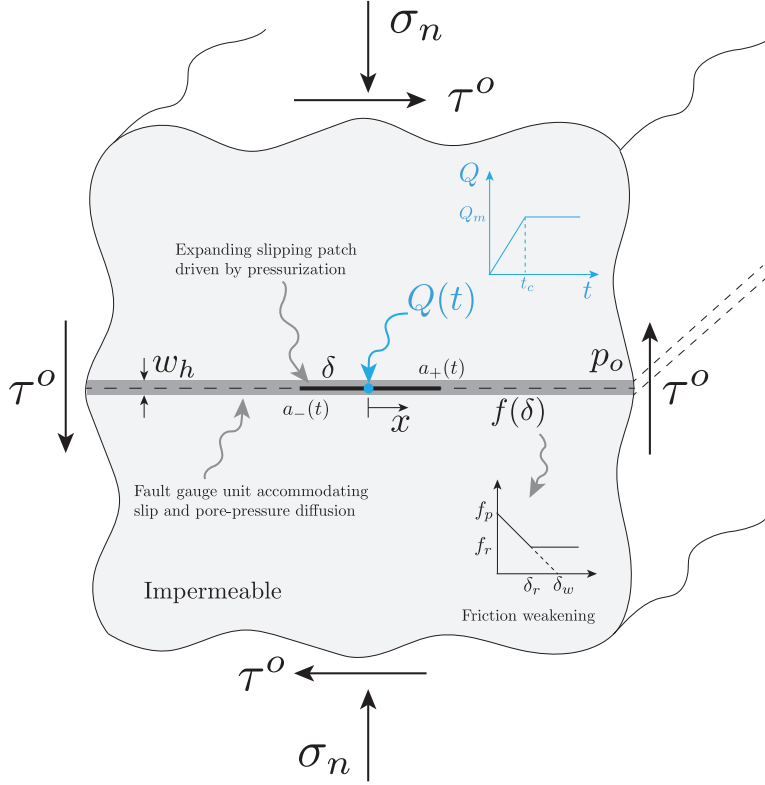


Figure 1: Plane-strain fault model subjected to a far-field stress state and fluid injection. The conductive planar fault is embedded into an homogeneous, isotropic and linear-elastic medium characterized by a negligible hydraulic diffusivity. The thin fault gouge unit, therefore, accommodates accelerating slip due to friction weakening condition and pore-pressure diffusion.

110 Prior fluid injection, we assume that the fault is in static equilibrium with the
 111 uniform in-situ stress state (locked status). This tacitly assumes that there
 112 is no effect of remote plates loading that would cause steady movements
 113 with slow energy release (creep) (Chen and Bürgmann, 2017). The ambient
 114 equilibrium is violated throughout pressurization, during which pore-pressure
 115 perturbation diffuses along the fault plane and activates a symmetric shear
 116 crack of length $2a$. This scenario of mechanically weak, thin fault gouge
 117 unit accommodating slip and pore fluid flow may be representative of an
 118 immature deep fault, whose hydraulic conductivity is much larger than the
 119 relatively undamaged rock around it (and thus impermeable host medium is

120 a reasonable assumption).

121 Before presenting the governing equations, it must be noted that we as-
 122 sume normal stresses positive in compression and shear stresses positive for
 123 clockwise rotation. Furthermore, thermal effects, dilatancy/compaction dur-
 124 ing shear deformations and poroelastic stress changes in the surrounding
 125 medium are neglected.

126 *2.1. Governing equations for quasi-static slip development driven by pore-*
 127 *pressure diffusion*

128 *2.1.1. Static equilibrium and constitutive law for frictional slip*

129 We consider the activation and propagation of a symmetric shear crack
 130 of length $2a$ driven by pore fluid flow inside the conductive fault plane. The
 131 shear stress τ at a given time t and position x on the slip surface is linearly
 132 related to slip δ (or shear displacement discontinuity) within the slipping
 133 region via the following quasi-static elastic equilibrium equation

$$\tau(x, t) = \tau^o - \frac{E_p}{4\pi} \int_{a_-(t)}^{a_+(t)} \frac{\partial\delta(\zeta, t)/\partial\zeta}{\zeta - x} d\zeta, \quad (1)$$

134 where a_+ and a_- are respectively the positive and negative positions of the
 135 crack tips, τ^o is the uniform background shear stress, $E_p = \frac{2G}{1-\nu}$ is the plane-
 136 strain Young's modulus (with G and ν being shear modulus and Poisson's
 137 ratio, respectively), $\partial\delta(\zeta, t)/\partial\zeta$ is the shear dislocation density along the
 138 crack that must satisfy the following condition

$$\lim_{\zeta \rightarrow a_{\pm}(t)} \frac{\partial\delta(\zeta, t)}{\partial\zeta} \sqrt{a \mp \zeta} = 0 \quad (2)$$

139 in order to remove stress singularity at crack tips (Uenishi and Rice, 2003),
 140 and $\frac{1}{\zeta - x}$ is the non-local elastic kernel. For merely planar frictional prob-
 141 lems that do not account for dilatancy or compaction during crack propa-
 142 gation, the elastic kernel affects only the shear stress distribution along the
 143 fault plane, while the uniform total normal stress σ_n remains constant.

144 Inside the sliding region, the force balance requires that the shear stresses
 145 must equal the available frictional resistance which we assume here to obey
 146 the Mohr-Coulomb yielding criterion (without cohesion), accounting for a
 147 slip weakening of friction coefficient

$$\tau(x, t) = f(\delta) (\sigma_n - p(x, t)), \quad |x| \leq a(t) \quad (3)$$

148 where $(\sigma_n - p(x, t)) = \sigma'_n(x, t)$ is the local effective normal stress, strictly
 149 function of pore-pressure evolution $p(x, t)$ inside the fault, and $f(\delta)$ is the
 150 slip weakening friction coefficient

$$f(\delta) = \begin{cases} f_p - \frac{(f_p - f_r)}{\delta_r} \cdot \delta & \delta \leq \delta_r \\ f_r & \delta > \delta_r \end{cases} \quad (4)$$

151 The frictional resistance, therefore, weakens linearly with shear deformations
 152 δ , from a peak value associated with peak friction coefficient f_p , to a
 153 residual value, after sufficiently large slip δ_r (such that the friction coefficient
 154 drops to its residual value f_r). Although more complicated frictional laws can
 155 be considered in (3), we use here the simple piece-wise linear weakening of
 156 friction coefficient with slip, which can be shown to be a good approximation
 157 of the phenomenological rate- and state- friction law at large slip rates and
 158 for small values of a/b (Uenishi and Rice, 2003; Rubín and Ampuero, 2005),
 159 or when $\Delta f_p/b \gg 1$, with Δf_p being the peak change of friction coefficient
 160 from the steady-state value (Garagash, 2021).

161 2.1.2. Pore-pressure diffusion

162 Under the assumption of negligible fluid leak-off in the surrounding elastic
 163 medium, fluid flow is confined within the fault gouge unit characterised by a
 164 constant hydraulic aperture w_h . Upon injection of volumetric flow rate $Q(t)$
 165 at $x = 0$, the diffusion of pore fluid over-pressure $\bar{p}(x, t) = p(x, t) - p_o$ is
 166 governed by the width-averaged fluid mass conservation equation

$$w_h c_f \frac{\partial \bar{p}}{\partial t} + \frac{\partial w_h v}{\partial x} = 0, \quad (5)$$

167 where $c_f [M^{-1}TL^2]$ is a parameter that combines pore fluid compressibility
 168 and pore space expansivity and v is the gap-averaged fluid flow velocity given
 169 by the Poiseuille's law

$$v = -\frac{k_f}{\mu} \frac{\partial \bar{p}}{\partial x}, \quad (6)$$

170 where $k_f [L^2]$ is the longitudinal fault permeability and $\mu = 12\mu' [MT^{-1}L^{-1}]$
 171 is a viscosity parameter (with μ' as dynamic viscosity of the fluid).

172 The point injection boundary condition requires that

$$w_h v = \pm \frac{Q(t)}{2}, \quad \text{at } x = 0^\pm \quad (7)$$

173 where the volumetric flow rate $Q(t)$ [L^2T^{-1}] increases linearly with time t ,
 174 up to reach a plateau after a given time t_c , i.e.

$$Q(t) = \begin{cases} \frac{Q_m}{t_c} \cdot t & t \leq t_c \\ Q_m & t > t_c \end{cases} \quad (8)$$

175 Since we neglect dilatation/compaction of fault gouge unit during shear
 176 crack propagation, fault permeability k_f and hydraulic aperture w_h remain
 177 constant throughout pressurization. Equations (5) and (6), therefore, reduce
 178 to the well-known parabolic diffusion equation

$$\frac{\partial \bar{p}}{\partial t} - \alpha \frac{\partial^2 \bar{p}}{\partial x^2} = 0 \quad (9)$$

179 that govern over-pressure diffusion inside the fault conduit characterized by
 180 a constant hydraulic diffusivity $\alpha = \frac{k_f}{c_f \cdot \mu}$ [L^2T^{-1}]. Using specific boundary
 181 and initial conditions representative of the particular injection scenario and
 182 its time history, equation (9) can be solved analytically for the spatial and
 183 temporal evolution of pore-fluid over-pressure along the fault plane $\bar{p}(x, t)$
 184 (see Appendix A for full details). During the ramp-up phase of injection
 185 rate, i.e. for $t \leq t_c$, the over-pressure evolution in function of time t and the
 186 normalized coordinate $\xi = \frac{x}{\sqrt{4\alpha t}}$ is given by

$$\bar{p}(\xi, t \leq t_c) = \underbrace{\left(\frac{2Q_m \mu \sqrt{\alpha t^{3/2}}}{3k_f \sqrt{\pi} t_c w_h} \right)}_{\Delta P(t)} \cdot \underbrace{\left(e^{-\xi^2} (1 + \xi^2) - \sqrt{\pi} |\xi| \left(\frac{3}{2} + \xi^2 \right) \text{Erfc}(|\xi|) \right)}_{\Psi(\xi)}, \quad (10)$$

187 where Erfc is the complementary error function. Notice that the analytical
 188 solution (10) is expressed as a product two independent functions, $\Delta P(t)$
 189 and $\Psi(\xi)$, which identify respectively the maximum over-pressure evolution
 190 at injection point and its instantaneous spatial distribution.

191 The pore-fluid evolution after the ramp-up phase, instead, is governed by the
 192 following equation (Cole et al., 2011)

$$\bar{p}(x, t > t_c) = \int_{-\infty}^{\infty} G(x-x', t-t_c) \cdot \bar{p} \left(\frac{x'}{\sqrt{4\alpha t_c}}, t_c \right) dx' + \frac{\alpha Q_m \mu}{w_k k_f} \int_{t'=t_c}^t G(x, t-t') dt' \quad (11)$$

193 where $\bar{p}\left(\frac{x'}{\sqrt{4\alpha t_c}}, t_c\right)$ denotes the over-pressure distribution at time t_c and
 194 $G(x - x', t - t')$ is the fundamental heat conduction solution valid for an
 195 infinite one-dimensional body subjected to an instantaneous point source
 196 (also called Green's function) (Carslaw and Jaeger, 1959). We revert the
 197 reader to Appendix A for its analytical expression. It is worth mentioning
 198 that the solution (11) is valid for $t > t_c$ and it takes into account the whole
 199 injection history during the ramp-up phase. If t_c vanishes, then we recover the
 200 analytical solution of pressurization at constant injection rate (see Appendix
 201 A).

202 In this contribution, we assume that maximum over-pressure occurring
 203 at injection point $x = 0$ remains always below the ambient effective normal
 204 stress $\sigma'_o = \sigma_n - p_o$ applied on the fault plane, i.e.

$$\frac{\bar{p}(0, t)}{\sigma'_o} < 1, \quad \forall t \quad (12)$$

205 implying that the minimum principal effective stresses $\sigma'_n(x, t)$ remain always
 206 compressive (positive) throughout pressurization and hydraulic fracturing
 207 type of failure never occurs (which would require the full coupling between
 208 flow and elastic deformations).

209 The along-fault pore-pressure diffusion changes the local effective normal
 210 stresses and hence drives the symmetric slip propagation when the Mohr-
 211 Coulomb criterion (3) is locally violated. Shear deformations, instead, do not
 212 affect pore-pressure evolution since shear-induced dilatancy or compaction
 213 is neglected and so fault permeability / porosity changes too. Although
 214 this assumption is debatable, since such inelastic deformations do affect slip
 215 stability on a planar fault with frictional weakening properties (Garagash
 216 and Rudnicki, 2003; Zhang et al., 2005; Ciardo and Lecampion, 2019), we
 217 want to minimise the complexities in the model and focus solely on the effect
 218 of injection rate ramp-up on the potential nucleation and arrest of dynamic
 219 slip. The hydro-mechanical model, therefore, is only one-way coupled and
 220 it is equivalent to the one proposed by Garagash and Germanovich (2012),
 221 with the difference that the point injection volumetric rate is not constant
 222 in time but changes according to (8).

223 We have introduced in the model an additional parameter t_c , therefore
 224 we expect another dimensionless parameter governing the hydro-mechanical
 225 fault's response (on top of the ones introduced by Garagash and Germanovich
 226 (2012)). For sake of completeness we present and discuss in the next section

227 all the dimensionless governing parameters resulting from scaling analysis
228 (see Appendix B for more details), as well realistic values that are then used
229 in the numerical simulations.

230 2.2. Dimensionless governing parameters

231 Upon normalization of all the governing equations (1-8) following Uenishi
232 and Rice (2003); Garagash and Germanovich (2012) (see Appendix B), the
233 hydro-mechanical fault response depends only on four dimensionless param-
234 eters:

- 235 • Stress criticality $\frac{\tau^o}{\tau_p}$, which represents the closeness of the ambient fault
236 stress state to failure (and thus to the peak shear strength $\tau_p = f_p \sigma'_o$).
237 Levandowski et al. (2018) claims that stress criticality is the most im-
238 portant factor for induced earthquakes hazard. Favourably oriented
239 frictional weakening faults with respect to the in-situ stress field, typ-
240 ically characterised by a large stress criticality ($\frac{\tau^o}{\tau_p} \lesssim 1$), are very
241 susceptible to host run-away seismic ruptures. Indeed, a little stress
242 perturbation is sufficient to re-activate slip and its velocity propaga-
243 tion tends to diverge rapidly due to friction weakening and possibly
244 other weakening mechanisms, such us flash-heating and thermal pres-
245 surization (Viesca and Garagash, 2015). Garagash and Germanovich
246 (2012) have shown via a stability analysis that critically stressed pres-
247 surized faults, for which the relation $\tau^o > \tau_r = f_r \sigma'_o$ is strictly satis-
248 fied, host always the nucleation of an unabated dynamic event. Such
249 a run-away rupture, however, can be suppressed when shear-induced
250 dilatancy kicks-off and dilatant hardening stabilises slip propagation
251 (Lockner and Byerlee, 1994; Segall et al., 2008; Ciardo and Lecam-
252 pion, 2019). Critically stressed faults in seismogenic zones have been
253 observed in Oklahoma and Southern Kansas (Qin et al., 2019), in the
254 German continental deep drillhole (KTB) (Ito and Zoback, 2000) and
255 in central California along the San Andreas fault system (Zoback et al.,
256 1987; Rice, 1992) to cite a few examples.

257 On the other hand, fault zones not favourably oriented to the local
258 in-situ stress field are characterised by low stress criticality ($\frac{\tau^o}{\tau_p} \gtrsim 0$),
259 which implies that larger over-pressures are required to activate slip. By
260 making the analogy with critically stressed pressurized faults exhibiting

261 linear slip-weakening behaviour, low stress criticality and $\tau^o < \tau_r$ im-
 262 plies a quasi-static, stable slip propagation with eventually a nucleation
 263 and arrest of a dynamic event at large pressurization time (Garagash
 264 and Germanovich, 2012).

265 The role of the initial effective stress state on the nature of seismicity
 266 has been also investigated in fault lab experiments by Passelègue et al.
 267 (2020), where they show that faults with similar frictional properties
 268 can rupture at both slow and fast rupture velocity depending on their
 269 initial stress criticality, in agreement with theoretical predictions based
 270 on linear elastic fracture mechanics (LEFM).

271 • Friction weakening ratio $\frac{f_r}{f_p} = 1 - \frac{\delta_r}{\delta_w}$ that governs the shear stress
 272 drop within the crack tips (with $\delta_w = \frac{f_p}{(f_p - f_r)}\delta_r$ being the amount of
 273 slip at which the friction coefficient goes to zero if an unlimited linear
 274 slip-weakening friction law is considered). According to slip laboratory
 275 experiments of granite intact specimens, the slip weakening distance
 276 δ_r is approximately half a millimetre (Rice, 1980), but it can drop
 277 even below 0.1 mm (Wong, 1986). Similarly, δ_w is on the order of
 278 fault’s asperities and thus ranges between 0.1 and 10mm. The friction
 279 weakening ratio, therefore, can assume any values within the interval
 280 $(0, 1)$.

281 • Normalized maximum injection rate $\frac{Q_m}{Q^*}$, where $Q^* = \frac{2\sigma'_o w_h k_f}{a_w \mu}$ is the
 282 characteristic injection rate scale and $a_w = \frac{E_p}{2\tau_p}\delta_w$ is the slipping patch
 283 length-scale. Its value is typically in the order of a meter, but roughly
 284 one order of magnitude of variation is plausible due to the variation
 285 of δ_w with fault’s roughness and σ'_o with hydrogeological conditions.
 286 In normally pressurized formations, typically located in the Earth’s
 287 upper crust and characterised by lithostatic gradient and hydro-static
 288 pore-pressure conditions (Brace and Kohlstedt, 1980; Grawinkel and
 289 Stockhert, 1997), σ'_o may be a fraction of megapascals, while a boost
 290 to hundred of megapascals may be obtained in over-pressurized forma-
 291 tions located below the fluid retention depth (Suppe, 2014). Assuming
 292 a fluid viscosity parameter $\mu \sim 10^{-3}\text{Pa} \cdot \text{s}$ (water), a fault gouge per-
 293 meability k_f in the order of $\sim 10^{-16}\text{m}^2$ (Wibberley et al., 2008) and

hydraulic width w_h of few millimetres, the injection rate characteristic scale Q^* ranges between $\sim 10^{-9}\text{m}^2/\text{s}$ and $10^{-6}\text{m}^2/\text{s}$. If we assume that the out-of-plane fault length is ~ 1 km, then the volumetric injection rate scale would range between a fraction to few litres per second.

The maximum injection rate Q_m is a design parameter that can vary considerably from tens or hundreds litres per second in hydraulic-fracturing operations (see (Holland, 2013) for one example where > 160 l/s were injected in south-central Oklahoma), to few or fractions of litres per second for hydro-shearing stimulations (see for instance hydro-shearing experiments in Grimsel Test Site, Switzerland, where the maximum injection rate never exceeded 0.5 l/s (Villiger et al., 2020)).

The dimensionless parameter $\frac{Q_m}{Q^*}$, therefore, can assume relatively low or large values depending on the specific problem configuration. Since hydraulic fracturing type of failure is not considered in our model, all the numerical results that will be presented later are characterised by a relatively low/moderate value of $\frac{Q_m}{Q^*}$.

- $\frac{t_c}{t_w}$ \rightarrow ratio between the ramp-up time scale and along-fault diffusion

time-scale $t_w = \frac{a_w^2}{4\alpha}$. This ratio may vary considerably due to the variation of slipping patch length-scale a_w previously discussed, the specific fault hydraulic properties considered as well as the choice of the parameter t_c . Assuming a compressibility parameter $c_f \sim 0.5 \text{ GPa}^{-1}$ (Wibberley, 2002) and the hydraulic parameters previously defined (albeit the fault longitudinal permeability may range within the interval $10^{-19} - 10^{-15} \text{ m}^2$), we estimate an hydraulic diffusivity of $\alpha \sim 2 \cdot 10^{-4} \text{ m}^2/\text{s}$, which leads to a diffusion time scale of $t_w \sim 21$ min for $a_w = 1$ m.

The parameter t_c may be in the order of several minutes for relatively fast fault pressurizations (see for instance one cycle of the injection protocol used for one hydro-shearing stimulation in Grimsel, Switzerland (Amann et al., 2018)), but it may get up to several hours or few days for relatively slow pressurizations (see for instance the injection protocol used for hydraulic stimulation of Pohang Enhanced Geothermal System, 2017 (Yeo et al., 2020; Hofmann et al., 2018)). The ratio t_c/t_w , therefore, may be very large, in the order 1 or possibly very small.

328 In the following we explore the parameter space identified by these di-
329 mensionless ratios via numerical simulations. We vary them systematically
330 in order to investigate their impact on the nucleation and arrest (if occur) of
331 dynamic fault slip.

332 3. Numerical results

333 We use a fast boundary element based solver in order to solve numerically
334 the one-way coupled hydro-mechanical problem (1-11) (see details in (Ciardo
335 et al., 2020)). It is suited for 2D non-linear geo-mechanical problems involv-
336 ing localized inelastic deformations along pre-existing structural discontinu-
337 ities, such as faults or fractures. The elasto-static balance of momentum (1)
338 is discretized using displacement discontinuity method that, together with
339 the discretized form of shear weakening Mohr-Coulomb criterion (3), lead
340 to a non-linear system of equations for the unknowns inelastic deformations
341 (or displacement discontinuities) and effective tractions. For a given pore-
342 pressure history calculated using (10) and (11) (with the latter evaluated
343 numerically using a trapezoidal quadrature rule), such a resulting system is
344 solved iteratively using fixed point iterations combined with under-relaxation
345 (Quarteroni et al., 2000), and adopting a fully implicit integration scheme in
346 time. A speed up of computations is obtained via hierarchical approxima-
347 tion of the fully populated elasticity matrix (on top of memory reduction),
348 together with the use of an adequate block pre-conditioner that improves
349 spectral properties of the resulting matrix of coefficients (see full details in
350 (Ciardo et al., 2020)).

351 All the numerical results that will be presented in the following are obtained
352 by considering a sufficiently long planar fault such that $L_o/a_w = 40$, with
353 L_o being its half-length. The planar fault is then uniformly discretized with
354 $2 \cdot 10^3$ equal-sized straight elements, resulting in having 50 elements inside
355 the slip weakening region near crack tips (sufficient to accurately capture the
356 model non-linearity).

357 3.1. Ultimately stable fault ($\tau^o < \tau_r$)

358 Firstly, we present the case of fluid injection into an ultimately stable fault
359 characterised by $\tau^o/\tau_r < 1$, where $\tau_r = f_r \sigma'_o$ is the residual frictional strength
360 prior pressurization. This condition automatically implies that the fault is far
361 from being critically stressed, and thus the ratio τ^o/τ_p is kept relatively low.
362 By fixing the frictional weakening ratio to $f_r/f_p = 0.6$, we investigate the

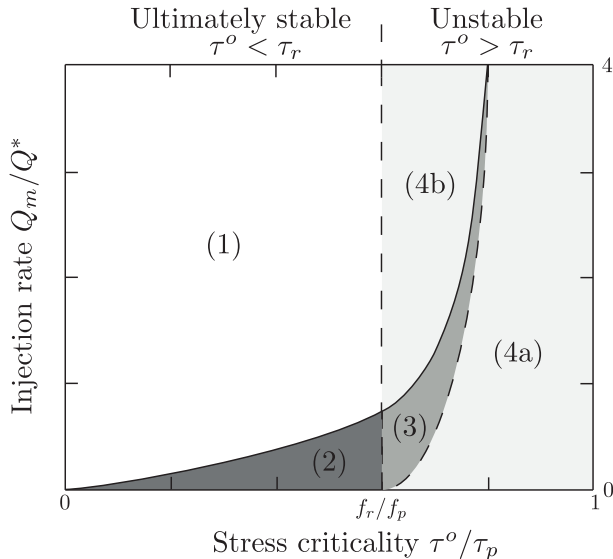


Figure 2: Map of slip regimes for a frictional weakening pressurized fault (adapted from Garagash and Germanovich (2012)), as function of stress criticality τ^o/τ_p , friction weakening ratio f_r/f_p and normalized (constant) injection rate Q_m/Q^* . Region (1) corresponds to the scenario of an ultimately stable fault in which pressurization leads always to quasi-static (stable) slip propagation. Region (2) corresponds to the case of quasi-static slip propagation, followed by a nucleation and arrest of dynamic slip. Regions (4b) and (4a), instead, represent the scenarios of an unstable fault exhibiting quick quasi-static crack propagation followed by an unabated dynamic rupture, whose nucleation is affected or not by residual friction coefficient f_r , respectively. Finally, in region (3) the nucleation of a run-away rupture on an unstable fault is preceded by a finite-sized seismic event.

363 fault's hydro-mechanical response for different values of normalized maximum
 364 injection rate Q_m/Q^* , in the two plausible limiting scenarios of $t_c/t_w \ll 1$
 365 and $t_c/t_w \gg 1$.

366 3.1.1. Fast ramp-up: the limit when $t_c/t_w \ll 1$

367 If the diffusion time-scale t_w is much larger than the ramp-up time t_c
 368 of injection rate, slip activation and propagation occurs when the injection
 369 rate has already reached its maximum value Q_m/Q^* . The quick ramp-up
 370 of injection rate has a negligible effect on pore-pressure evolution and the
 371 hydro-mechanical fault response is essentially driven by injection at constant
 372 volumetric rate. Garagash and Germanovich (2012) studied extensively the
 373 effect of such a type of fault pressurization on the nucleation and poten-
 374 tial arrest of a dynamic event. They came up with a slip regimes diagram

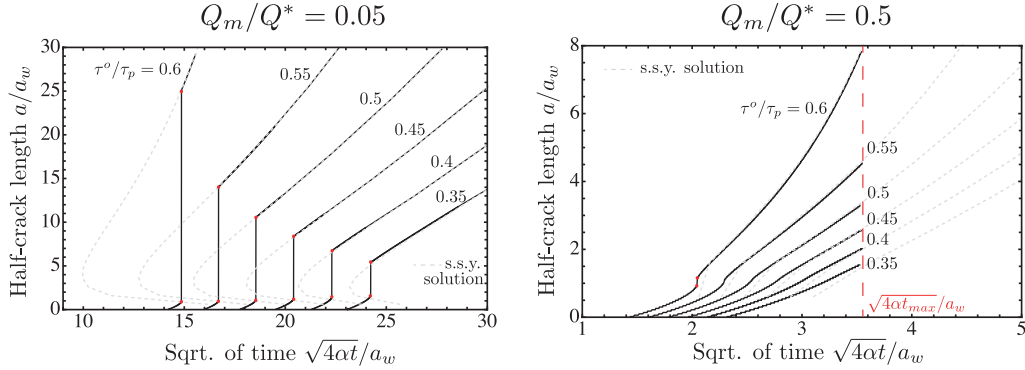


Figure 3: Time evolution of normalized half-crack length a/a_w (black solid lines) as function of various stress criticality values τ^o/τ_p , for two values of maximum injection rates $Q_m/Q^* = 0.05, 0.5$ and $t_c/t_w = 0.01 \ll 1$. The friction weakening ratio is set to $f_r/f_p = 0.6$, resulting in ultimately stable conditions, i.e. $\tau^o < \tau_r$ for each value of ambient fault loading condition. Grey dashed lines represent the small scale yielding (s.s.y.) solutions (see Appendix C) associated with constant injection rate type of pressurization, which represents a good approximation of the hydro-mechanical fault response when $t_c/t_w \ll 1$. In this particular example, indeed, the normalized ramp-up time is $\sqrt{4\alpha t_c}/a_w = 0.1$ and thus shear crack activation and propagation is essentially driven at constant injection rate. Red dots, instead, denote the nucleation and arrest of dynamic slip.

375 reported in Figure 2, in which a frictional weakening planar fault may experi-
 376 ence a finite-sized dynamic event (region (2)), a run-away rupture (affected
 377 or not by the residual friction coefficient f_r , region (4b) and (4a) respec-
 378 tively), a finite-sized dynamic event followed by an unabated seismic rupture
 379 (region (3)), or only aseismic slip (region (1)). This depends on the particu-
 380 lar set of dimensionless parameters that identifies a specific initial fault
 381 loading condition, pressurization and frictional property. Their theoretical
 382 and semi-analytical results, therefore, provide benchmark solutions for our
 383 numerical results.

384 Under ultimately stable conditions (i.e. $\tau^o < \tau_r$, left side of Figure 2), Gara-
 385 gash and Germanovich (2012) proved via a stability analysis that a fault ex-
 386periences a stable, quasi-static growth of slipping patch throughout sustained
 387 pressurization (i.e. aseismic slip). However, for sufficiently low injection rates
 388 Q_m/Q^* , the fault exhibits a finite-sized dynamic event, whose nucleation
 389 time t_n is considerably larger than fluid diffusion time-scale t_w (and thus
 390 $t_c \ll t_w \ll t_n$). These considerations are also confirmed with our numeri-

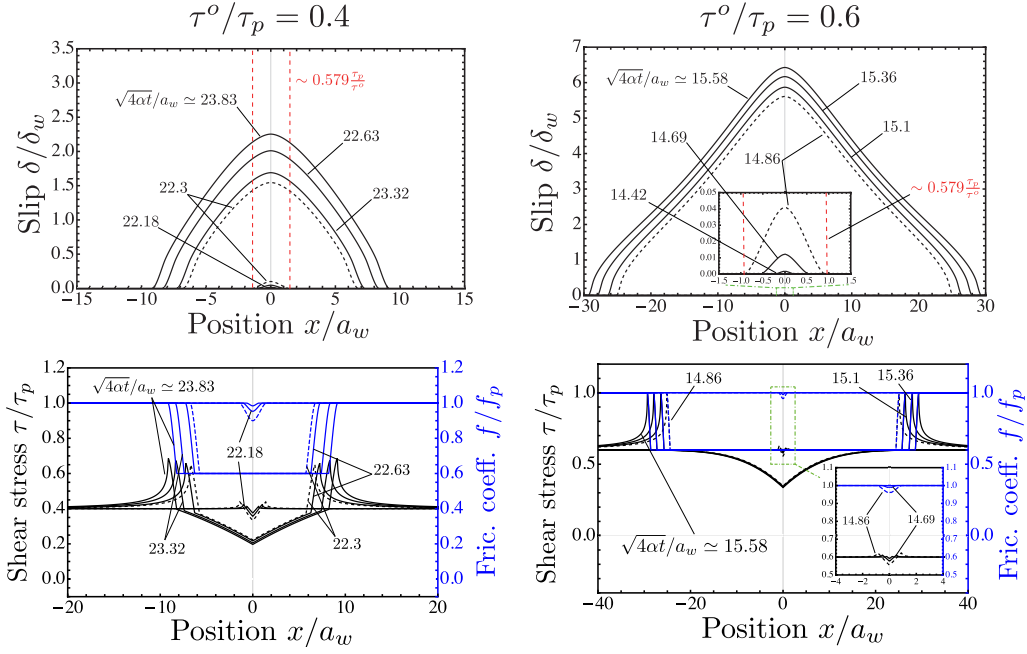


Figure 4: Normalized slip δ/δ_w , shear stress τ/τ_p and friction coefficient f/f_p profiles at different normalized time snapshots, for the case of maximum injection rate $\frac{Q_m}{Q^*} = 0.05$, ramp-up time scale $t_c/t_w = 0.01 \ll 1$ and two stress criticality values, i.e. $\tau^o/\tau_p = 0.4$ and 0.6 (ultimately stable faults as $f_r/f_p = 0.6$). Dashed blue and black lines correspond to the numerical solutions at nucleation time t_n (or in normalized form $\frac{\sqrt{4\alpha t_n}}{a_w}$), while dashed vertical red lines denote the asymptotic solution (13) provided by Garagash and Germanovich (2012).

391 cal results reported in Figure 3, where the normalized time evolution of half
 392 crack length a/a_w is displayed for two different values of $Q_m/Q^* = 0.05, 0.5$
 393 and different values of low/moderate stress criticality τ^o/τ_p (with the ratio
 394 t_c/t_w equal to 0.01 and τ^o always below or equal τ_r).

395 For the case of low injection rate $Q_m/Q^* = 0.05$, a nucleation of dynamic
 396 event followed by an arrest always occurs for each value of stress criticality
 397 considered, with larger dynamic run-out distances for increasing values of
 398 τ^o/τ_p (see Figure 3-left). Since the ambient effective stress states applied on
 399 the fault plane are far from failure, resulting in a stable quasi-static slipping
 400 patch propagation before reaching the nucleation time t_n (or in normalized
 401 form $\sqrt{4\alpha t_n}/a_w$), the shear crack tips are located well within the pressurized

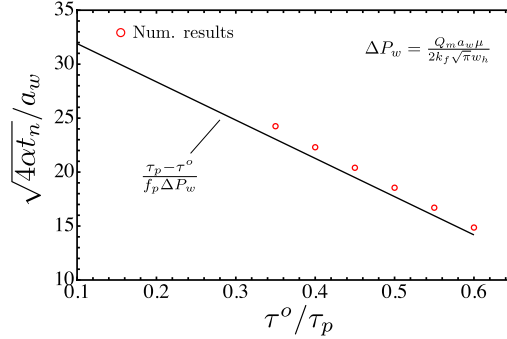


Figure 5: Comparison between the asymptotic solution (14) for ultimately stable faults valid for $t_c/t_w \ll 1$ and numerical solutions corresponding to the case of $t_c/t_w = 0.01$ and $Q_m/Q^* = 0.05$ (for which the necessary condition (15) is strictly satisfied for all values of τ^o/τ_p considered).

402 region. At nucleation time, the rupture is not affected by the residual friction
 403 coefficient f_r (due to the low slip accumulation during quasi-static slip
 404 propagation phase) and its half-length assumes the asymptotic expression
 405 (Garagash and Germanovich, 2012)

$$\frac{a_n}{a_w} \simeq 0.579 \cdot \frac{\bar{\tau}_p}{\tau^o} \quad (13)$$

406 A careful investigation of Figure 4 that displays the corresponding slip,
 407 shear stress and friction coefficient profiles at different time snapshots (with
 408 dashed lines corresponding to the ones at nucleation time) and for two stress
 409 criticality values $\tau^o/\tau_p = 0.4, 0.6$ confirms these theoretical predictions. It
 410 is worth noting at the arrest of the dynamic slip propagation, the friction
 411 coefficient has reached its residual value over almost the entire crack. The
 412 subsequent shear crack propagation, therefore, remains always stable in time
 413 as depicted in Figure 3-left. In addition to this, Garagash and Germanovich
 414 (2012) proposed an asymptotic solution for the nucleation time t_n that is valid
 415 for an ultimately stable fault that exhibits a dynamic event right after slip
 416 activation (similarly to the scenarios reported in Figure 3-left). Under this
 417 condition, indeed, the shear crack at nucleation time is confined near injection
 418 point and is subjected to a uniform over-pressure equal to the minimum value
 419 required to activate slip, i.e. $\bar{p}(x \simeq 0, t_n) = (\tau_p - \tau^o)/f_p$. This asymptotic

420 solution reads (Garagash and Germanovich, 2012)

$$\frac{\sqrt{4\alpha t_n}}{a_w} \simeq \frac{\tau_p - \tau^o}{f_p \Delta P_w}, \quad (14)$$

421 where $\Delta P_w = \frac{Q_m a_w \mu}{2k_f \sqrt{\pi} w_h}$ is the characteristic pore-pressure drop over the
 422 distance a_w . In order to guarantee that crack instability follows shortly after
 423 slip activation, ΔP_w must be very small compared to $\bar{p}(x \simeq 0, t_n)$, resulting
 424 in the following necessary condition

$$\frac{Q_m}{Q^*} \ll \sqrt{\pi} \left(1 - \frac{\tau^o}{\tau_p} \right) \quad (15)$$

425 The comparison between the nucleation times associated with numerical
 426 simulations reported in Figure 3-left (for which $Q_m/Q^* = 0.05$ strictly sat-
 427 isfies the condition (15) for each value of stress criticality considered) and
 428 the asymptotic solution (14) is displayed in Figure 5. A good agreement
 429 is obtained, specially for larger values of τ^o/τ_p for which the assumption of
 430 nucleation after activation is truly valid (see Figure 3-left). Obviously, lower
 431 values of Q_m/Q^* would lead to an earlier nucleation and thus a closer match
 432 between numerical and asymptotic solution.

433 For a larger value of injection rate $Q_m/Q^* = 0.5$ (such to fall into region
 434 (1) of Figure 2 for $\tau^o/\tau_p < 0.6$), instead, the slipping patch propagates al-
 435 ways quasi-statically for all values of τ^o/τ_p , before reaching the maximum
 436 pressurization time $\sqrt{4\alpha t_{max}}/a_w$ at which the condition (12) is violated (and
 437 at which all the simulations were stopped). One exception is a very small
 438 seismic event corresponding to the boundary case of $\tau^o/\tau_p = 0.6$ - see Figure
 439 3-right.

440 These numerical results confirm that in case the maximum flow rate Q_m
 441 is reached before pore fluid can actually diffuse into the fault, the hydro-
 442 mechanical fault response is essentially driven by pressurization at constant
 443 volumetric rate and a good match with theoretical results of Garagash and
 444 Germanovich (2012) is obtained. This is further strengthened by looking at
 445 the comparison between the numerical results reported in Figure 3 (black
 446 lines) and the small-scale yielding (s.s.y.) asymptotic solutions (dashed grey
 447 lines) associated with pressurization at constant injection rate only (see Ap-
 448 pendix C). For $a/a_w \gtrsim 2$, indeed, a perfect match is obtained, suggesting
 449 that the effect of the quick ramp-up phase on slip propagation is negligible.

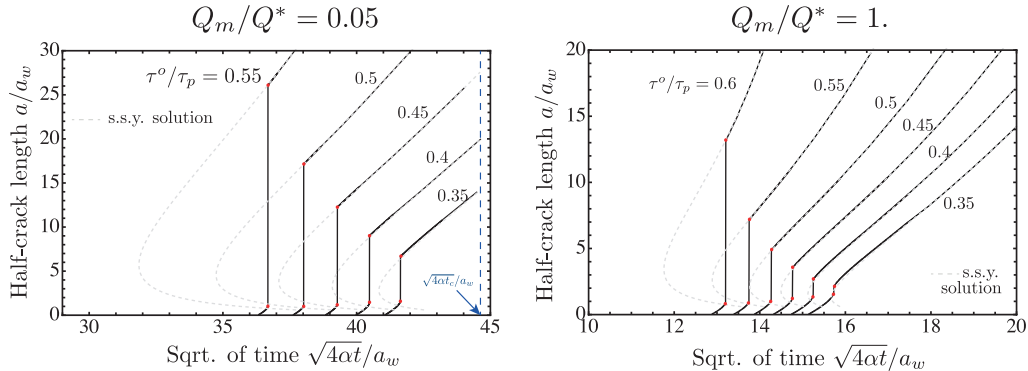


Figure 6: Time evolution of normalized half-crack length a/a_w (black solid lines) as function of stress criticality τ^o/τ_p , for two values of maximum injection rates $Q_m/Q^* = 0.05, 1$ and $t_c/t_w = 2 \cdot 10^3 \gg 1$. The friction weakening ratio is set to $f_r/f_p = 0.6$, resulting in ultimately stable conditions, i.e. $\tau^o < \tau_r$ for each value of ambient fault loading condition. Grey dashed lines represent the small scale yielding (s.s.y.) solution associated with linear ramp-up of injection rate (see Appendix C), whose normalized ending time in this example is $\sqrt{4\alpha t_c}/a_w = 44.72$. Red dots, instead, denote the nucleation and arrest of dynamic slip.

450 3.1.2. Slow ramp-up: the limit when $t_c/t_w \gg 1$

451 We present here the opposite scenario in which injection rate increases
 452 linearly in time but the maximum value is reached after the pore fluid sub-
 453 stantially diffuses along the fault plane (i.e. $t_c \gg t_w$). Under this condition,
 454 pore-pressure perturbation activates a shear crack (slip) during the ramp-up
 455 of injection rate, and the potential nucleation of a dynamic event would occur
 456 when pressurization is approaching the changing time t_c . In other words,
 457 $t_c \gg t_w$ and $t_n \lesssim t_c$.

458 Scaling analysis reported in Appendix B suggests that the pressurization
 459 rate, i.e. how quick is the ramp-up of injection rate before time t_c , does
 460 play a role in the hydro-mechanical fault response. In order to investigate its
 461 effect on the potential nucleation and arrest of dynamic slip on ultimately
 462 stable faults, we run several simulations keeping the ratio t_c/t_w constant (and
 463 much larger than 1) and varying the maximum injection rate Q_m/Q^* . This is
 464 equivalent to keeping the normalized injection rate constant and varying the
 465 (large) ratio t_c/t_w , resulting in relatively low or large pressurization rates.

466 Figure 6 displays the time evolution of half-crack length for different values
 467 of stress criticality, a ratio $t_c/t_w = 2 \cdot 10^3$ and two maximum injection rates
 468 $Q_m/Q^* = 0.05, 1$, representative of low and moderate fault pressurization

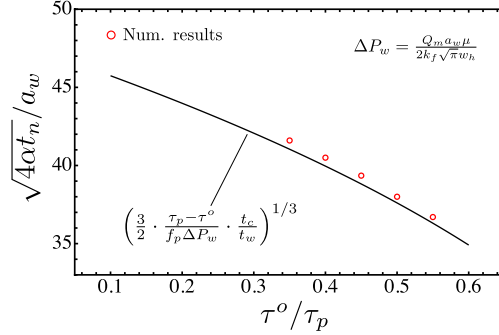


Figure 7: Comparison between the asymptotic analytical solution (16) for ultimately stable faults valid for $t_c/t_w \gg 1$ and numerical solutions corresponding to the case of $t_c/t_w = 2 \cdot 10^3$ and $Q_m/Q^* = 0.05$ (for which the necessary condition (15) is strictly satisfied for all values of τ^o/τ_p considered).

469 rates, respectively. Notice that the maximum injection rates considered are
 470 similar to the ones used in the previous case of $t_c/t_w \ll 1$ (see Figure 3),
 471 although the pressurization rates are considerably different. From a compar-
 472 ison of Figures 6-right and 3-right, we clearly observe that, in the case of
 473 $t_c/t_w \gg 1$, a moderate maximum injection rate is not sufficient to quench
 474 the finite-sized dynamic slip event for each value of stress criticality (unlike
 475 the case previously discussed). This is certainly due to the different type of
 476 fault pressurization that drives the slipping patch expansion. In this case,
 477 indeed, slip is driven by linear increase branch of injection rate and pore-
 478 pressure at injection point evolves proportional to $\sim t^{3/2}$ (see Eq. (10)). A
 479 further comparison of Figures 3-left and 6-left for the exact same value of
 480 $Q_m/Q^* = 0.05$ reveals that the dynamic run-out distances are rather similar
 481 for each corresponding value of stress criticality, while the nucleation times
 482 differ considerably. With the similar approach of Garagash and Germanovich
 483 (2012), we derive an asymptotic expression for the nucleation time that is
 484 valid when nucleation of dynamic slip follows shortly shear crack activation.
 485 By setting $\bar{p}(\xi \simeq 0, t_n)$ from equation (10) equal to $(\tau_p - \tau^o)/f_p$, i.e. equal
 486 to the minimum over-pressure required to activate slip, we get the following
 487 asymptotic expression for the normalized nucleation time

$$\frac{\sqrt{4\alpha t_n}}{a_w} \simeq \left(\frac{3}{2} \cdot \frac{\tau_p - \tau^o}{f_p \Delta P_w} \cdot \frac{t_c}{t_w}\right)^{1/3} \iff \frac{Q_m}{Q^*} \ll \sqrt{\pi} \left(1 - \frac{\tau^o}{\tau_p}\right), \quad (16)$$

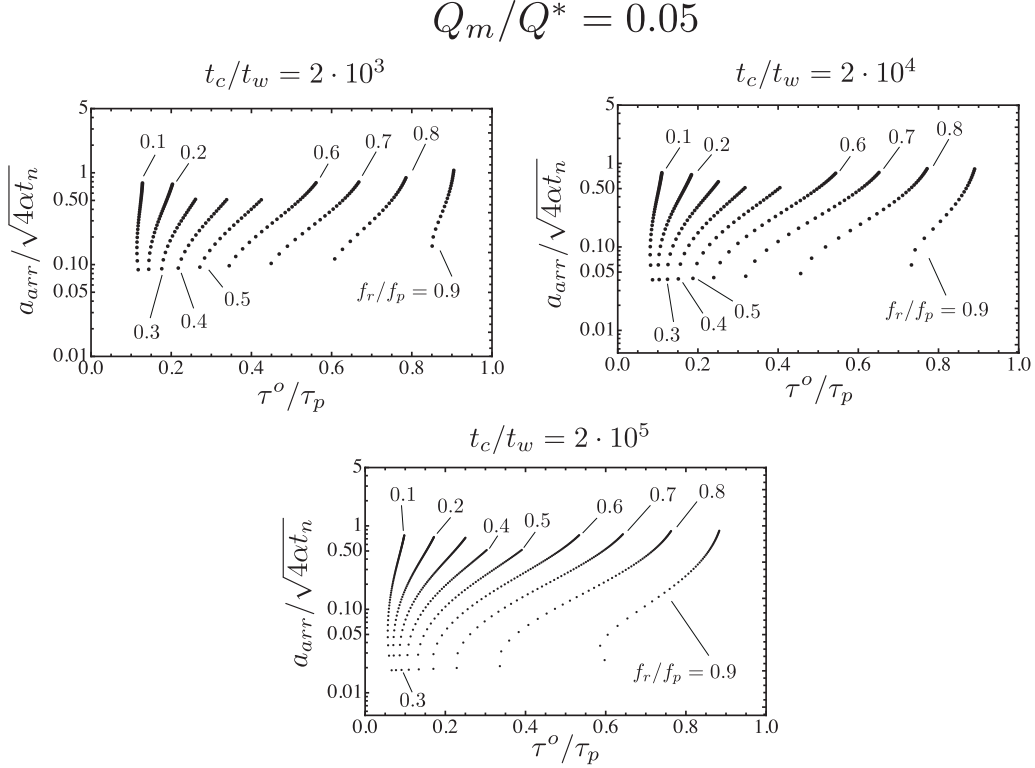


Figure 8: Normalized dynamic run-out distances $a_{arr}/\sqrt{4\alpha t_n}$ in function of stress criticality τ^o/τ_p and friction weakening ratio f_r/f_p , for different ramp-up scenarios of injection rate, i.e. different values of $\frac{Q_m t_w}{t_c Q^*}$. These latter are obtained by fixing the maximum injection rate at $Q_m/Q^* = 0.05$ and varying the (large) ratio t_c/t_w (notably $t_c/t_w = 2 \cdot 10^3, 2 \cdot 10^4, 2 \cdot 10^5 \gg 1$). In all the cases nucleation occurs prior reaching time t_c , even for low values of f_r/f_p .

488 where ΔP_w is the same characteristic pore-pressure drop reported in Eq.
 489 (14). Unlike the asymptotic solution (14) valid for $t_c/t_w \ll 1$, in this case the
 490 normalized nucleation time varies non-linearly with stress criticality τ^o/τ_p ,
 491 with a power law exponent equal to 1/3. Furthermore, it appears the depen-
 492 dency on the dimensionless ratio t_c/t_w with the same power law exponent
 493 (as expected from scaling analysis), revealing that the nucleation time of dy-
 494 namic slip does depend on how quick the injection rate ramp-up occurs in
 495 time, and thus on how large is the ratio $\frac{Q_m \cdot t_w}{Q^* \cdot t_c}$. The comparison between
 496 the numerical results corresponding to the case of $Q_m/Q^* = 0.05$, for which

497 the condition $\frac{Q_m}{Q^*} \ll \sqrt{\pi} \left(1 - \frac{\tau^o}{\tau_p}\right)$ is strictly satisfied for all the values of
 498 τ^o/τ_p , and the asymptotic solution (16) is displayed in Figure 7. A good
 499 agreement is obtained, with a closer match for larger values of τ^o/τ_p due to
 500 the earlier nucleations after slip activations (see Figure 6-left).

501

502 The asymptotic solution (16) gives also an estimation of the normalized
 503 position of the fluid front at nucleation time. A careful inspection of Figure
 504 6-left reveals that, at the onset of dynamic slip, the slipping patch front lags
 505 well within the pressurized region, and their relative distance decreases dra-
 506 matically after the arrest of dynamic event (with fluid front always located
 507 ahead the slip front). Furthermore, we can observe that our numerical results
 508 (black lines) match perfectly with the small-scale yielding (s.s.y.) asymptotic
 509 solutions associated with linear increase of injection rate (dashed grey lines -
 510 see Appendix C for more details) for each value of stress criticality. Since the
 511 extent of the arrested dynamic crack is always much larger than the slipping
 512 patch length scale a_w (thus s.s.y. condition is truly valid) and the corre-
 513 sponding nucleation time is known analytically from (16), we can calculate
 514 analytically the dynamic run-out distances a_{arr} directly from the small-scale
 515 yielding solution resolved at nucleation time. By simply replacing the nucle-
 516 ation time t_n obtained from Eq. (16) into the increment of stress intensity
 517 factor associated with ramp-up of injection rate (C.3), we can solve the re-
 518 sulting implicit equation obtained from propagation criterion (C.6) for the
 519 unknown arrested crack lengths. Obviously, this analytical equation is only
 520 valid in the case of early dynamic crack nucleation after activation, but it
 521 allows to investigate systematically the effect of the other dimensionless pa-
 522 rameters on the arrest of dynamic rupture.

523 We examine the effect of pressurization rate by fixing the maximum injection
 524 rate to $Q_m/Q^* = 0.05$ (such to satisfy the necessary condition in Eq. (16) for
 525 the whole range of stress criticality τ^o/τ_p) and varying the large ratio t_c/t_w ,
 526 obtaining thus a relatively quick or slow ramp-up of injection rate. In Figure
 527 8, we display the normalized dynamic run-out distances $a_{arr}/\sqrt{4\alpha t_n}$ in func-
 528 tion of stress criticality and friction weakening ratio f_r/f_p . We can observe
 529 that, for a given ramp-up of injection rate, the nucleation of a dynamic event
 530 is expected on a wider range of stress criticality for intermediate values of
 531 friction weakening ratio (compared to the cases of $f_r/f_p \rightarrow 0$ or $f_r/f_p \rightarrow 1$).
 532 Moreover, for a given value of f_r/f_p , the range of stress criticality in which
 533 a dynamic event is expected increases for increasing values of t_c/t_w (i.e. for

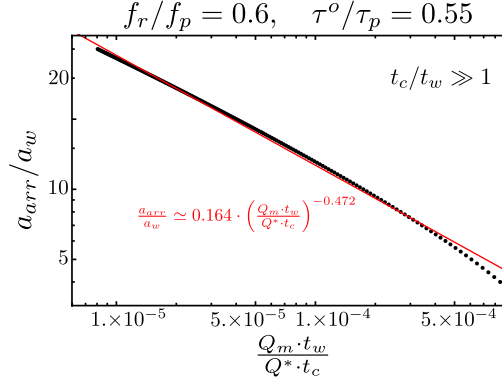


Figure 9: Normalized dynamic run-out distances a_{arr}/a_w (black dots) in function of pressurization rate $\frac{Q_m t_w}{t_c Q^*}$ (with $t_c/t_w \gg 1$), for an ultimately stable fault characterised by $\tau^o/\tau_p = 0.55$ and friction weakening ratio of $f_r/f_p = 0.6$. The red solid line represents a linear fit of the numerical results (black dots) in the log-log plot.

534 decreasing values of $\frac{Q_m t_w}{t_c Q^*}$, revealing that low pressurization rates promote
 535 the nucleation of a finite-sized dynamic rupture on ultimately stable faults.
 536 It is also interesting to note that the arrest of the dynamic rupture always
 537 occurs within the pressurized region, i.e. $\frac{a_{arr}}{\sqrt{4\alpha t_n}} \lesssim 1$, regardless of i) how
 538 quick the injection rate increase in time, ii) friction weakening ratio consid-
 539 ered and iii) stress criticality value. Furthermore, for a given value of f_r/f_p
 540 and τ^o/τ_p , the dynamic run-out distance increases for decreasing values of
 541 $\frac{Q_m t_w}{t_c Q^*}$, i.e. for slower ramp-up of injection rate. This can be grasped clearly
 542 from Figure 9 where the normalized ramp-up rate $\frac{Q_m t_w}{t_c Q^*}$ is plotted against
 543 the normalized dynamic run-out distance a_{arr}/a_w in a log-log plot. As one
 544 can see, the arrest of the dynamic slipping patch decreases non-linearly with
 545 increasing values of ramp-up rates, approximately proportional to an inverse
 546 power-law with exponent equal to -0.472 (see red line in Figure 9).

547 3.2. Unstable fault ($\tau^o > \tau_r$)

548 Finally, we present the case of injection into an unstable fault charac-
 549 terised by $\tau^o/\tau_r > 1$ at ambient conditions. In this particular scenario the
 550 fault is critically stressed and thus prompt to fail. A small strength pertur-
 551 bation, due to for instance a small pore-pressure increment \bar{p} , always leads to

552 a quick quasi-static shear crack propagation followed by a run-away dynamic
 553 rupture (albeit in some circumstances a finite-sized dynamic event may pre-
 554 ceede the run-away rupture - see region (3) of Figure 2 valid for the injection
 555 scenario at constant volumetric rate). Assuming a shear crack activation
 556 during the early ramp-up of injection rate (valid when t_c is not much smaller
 557 than t_w and thus pore fluid can diffuse within the fault plane), the slipping
 558 patch outpaces rapidly the pore fluid front and, at nucleation time t_n , the
 559 following condition hold

$$\frac{a_n}{a_w} \gg \frac{\sqrt{4\alpha t_n}}{a_w}, \quad (17)$$

560 with t_n much smaller than both fluid diffusion time-scale t_w and ramp-up
 561 time scale t_c . The pressurized region at nucleation time, therefore, is al-
 562 ways confined near injection point and the corresponding pore-pressure dis-
 563 tribution can be approximated using an equivalent point force distribution
 564 $\bar{p}(x, t) \simeq \Delta P(t) \delta_{dirac}(x)$, with $\Delta P(t)$ defined in (10).

565 Based on the previous work of Uenishi and Rice (2003), Garagash and Ger-
 566 manovich (2012) showed that there exist an asymptotic solution in terms of
 567 critical shear crack length that is universal (i.e. independent of the particular
 568 type of injection scenario) and it reads

$$\frac{a_n}{a_w} \simeq 0.579 \quad (18)$$

569 In addition to this, they developed semi-analytically an outer and inner
 570 asymptotic solutions that are valid at different fault position with respect
 571 to fluid front location (see Appendix D). The outer solution is so called
 572 because is valid outside the pressurization region, i.e. for $|x| \gg \sqrt{4\alpha t}$. The
 573 inner asymptotic solution, instead, is valid for $|x| \lesssim \sqrt{4\alpha t}$, i.e. near injection
 574 point due to the limited extension of pressurization region on unstable faults
 575 before instability.

576 Since the outer asymptotic solution is universal, i.e. it is valid for any type of
 577 peak pore-pressure distribution (see Appendix D), we can make use of such
 578 a solution to derive an asymptotic expression for the nucleation time t_n in
 579 function of the ramp-up of injection rate (and compare it with the solution of
 580 Garagash and Germanovich (2012) valid for pressurization at constant flow
 581 rate that can actually be retrieved here in the limit of $t_c \rightarrow 0$). Based on
 582 the outer solution, indeed, the integrated net over-pressure along the fault
 583 at crack instability is approximately given by (Garagash and Germanovich,

584 2012)

$$\Delta P(t_n) = \int_{-\infty}^{+\infty} \bar{p}(x, t_n) dx \simeq \mathcal{P} \frac{\tau_p - \tau^o}{f_p} a_w, \quad (19)$$

585 where $\mathcal{P} \simeq 0.8369$ is the scaled magnitude of the point force, independent of
 586 the type of fault pressurization. By replacing Eq. (10) into Eq. (19), after
 587 some algebra, we obtain

$$\frac{\sqrt{4\alpha t_n}}{a_w} \simeq \left(\frac{2^2 \cdot 0.8369}{\sqrt{\pi}} \cdot \frac{\tau_p - \tau^o}{f_p \Delta P_w} \cdot \frac{t_c}{t_w} \right)^{1/4}, \quad (20)$$

588 where ΔP_w is the same characteristic pore-pressure drop of Eq. (14). It is
 589 interesting to note that the normalized nucleation time in the case of an un-
 590 stable fault subjected to a ramp-up of injection rate varies non-linearly with
 591 both stress criticality τ^o/τ_p and pressurization rate $\frac{Q_m \cdot t_w}{Q^* \cdot t_c}$ (with a power
 592 law exponent equal to 1/4). A comparison with the normalized nucleation
 593 time associated with constant injection rate type of pressurization (Garagash
 594 and Germanovich, 2012)

$$\frac{\sqrt{4\alpha t_n}}{a_w} \simeq \left(\frac{2 \cdot 0.8369}{\sqrt{\pi}} \frac{\tau_p - \tau^o}{f_p \Delta P_w} \right)^{1/2} \quad (21)$$

595 suggests that the ramp-up of injection rate prior the maximum plateau may
 596 promote an earlier or later nucleation of a run-away dynamic event, with
 597 respect to the injection scenario at constant volumetric rate. Indeed, by
 598 taking the ratio between equation (20) and equation (21), we obtain $1.20636 \cdot$
 599 $\left(\frac{f_p \Delta P_w t_c}{\tau_p - \tau^o t_w} \right)^{1/4}$, implying that an earlier nucleation is expected when

$$\frac{t_c}{t_w} < \frac{1}{1.20636^4} \frac{\tau_p - \tau^o}{f_p \Delta P_w}, \quad (22)$$

600 for a given maximum value of injection rate Q_m .

601 In order to verify all these theoretical predictions, we run several numer-
 602 ical simulations with different stress criticality values, such to obtain highly
 603 critically stressed fault conditions for which condition (17) is strictly satisfied.
 604 We fixed the maximum injection rate at $Q_m/Q^* = 100$, the friction weaken-
 605 ing ratio at $f_r/f_p = 0.6$ and the normalized changing time at $t_c/t_w = 0.5$. As
 606 we can observe from Figure 10, the quick quasi-static shear crack propagation

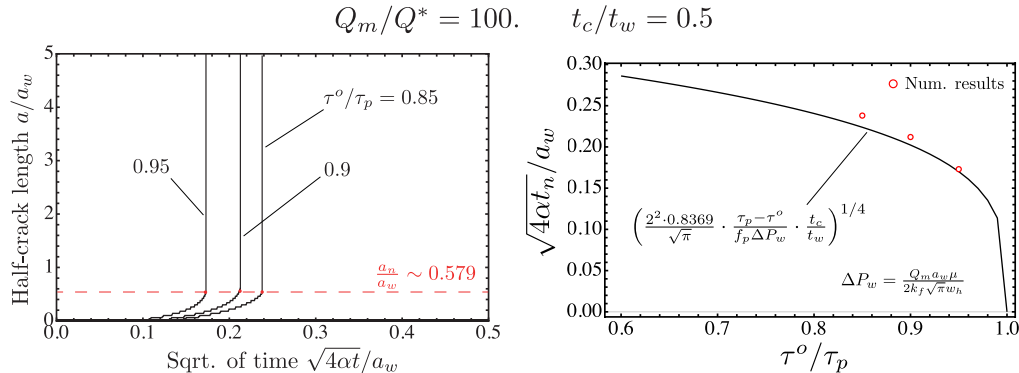


Figure 10: Left: time evolution of normalized half-crack length a/a_w for different values of (large) stress criticality $\tau^o/\tau_p = 0.95, 0.9, 0.85$. Since the friction weakening ratio is $f_r/f_p = 0.6$, all the scenarios correspond to very unstable faults. The maximum injection rate is set to $Q_m/Q^* = 100.$, while the normalized ramp-up time is $t_c/t_w = 0.5$ (or in terms of normalized square root of time $\sqrt{4\alpha t_c}/a_w \simeq 0.71$). Right: comparison between the asymptotic analytical solution (20) and the numerical results in terms of normalized nucleation time $\sqrt{4\alpha t_n}/a_w$, for each value of τ^o/τ_p considered (see left plot).

607 is always followed by a run-away dynamic rupture, whose nucleation time is
 608 in good agreement with theoretical prediction of Eq. (20) (with a better
 609 accuracy for larger values of τ^o/τ_p due to a more strict validity of condition
 610 (17)). At the instability, the normalized crack length a/a_w is approximately
 611 ~ 0.579 for each value of stress criticality considered (as expected) and the
 612 corresponding scaled slip distributions match very well the outer and inner
 613 asymptotic solutions at $|x| \gg \sqrt{4\alpha t}$ and $|x| \lesssim \sqrt{4\alpha t}$, respectively (see Fig-
 614 ures D.12 and D.13 in Appendix D).

615 4. Discussions

616 In this section, we first discuss the limitations of our modelling approach
 617 and their effects on nucleation and arrest of dynamic fault slip. After that,
 618 we discuss the implications of our results on injection-induced seismicity, a
 619 topic that has raised the attention of scientific community and public opinion
 620 since second half of 20th century (Simpson, 1986).

621 4.1. Model limitations

622 The most severe approximation in our model is the neglect of dilatancy
 623 during fluid-driven shear crack propagation. The assumption of impermeable

624 host medium around fault gouge unit accommodating slip and pore-fluid dif-
 625 fusion is representative of an immature deep fault, typically characterised by
 626 rough inner surfaces. Strong dilatant behaviour associated with sliding over
 627 fault’s asperities has been observed in laboratory experiments (Lockner and
 628 Byerlee, 1994; Samuelson et al., 2009) as well as during field experiments in
 629 the context of geothermal energy exploitation (Batchelor, 1985). Under the
 630 assumption of no fluid leak-off in the relative undamaged rock around the
 631 fault, shear-induced dilatancy does affect pore-pressure evolution, which in
 632 turn leads to a feedback on slip propagation (due to full coupling between
 633 flow and elastic deformations). Under undrained conditions, indeed, shear-
 634 induced dilatancy leads to a pore-pressure drop and thus to a local increase
 635 of effective normal stress (dilatant hardening) (Rudnicki, 1979; Segall and
 636 Rice, 1995). Ciardo and Lecampion (2019) have shown that such a dila-
 637 tant hardening effect, quantified by a scaled undrained pore-pressure drop
 638 $\frac{\Delta w_h}{w_h c_f \sigma'_o}$ with Δw_h being the increment of fault opening, does impact the
 639 transition between aseismic and seismic slip on frictional weakening imma-
 640 ture faults. They showed, in fact, that a large dilatant behaviour suppresses
 641 the nucleation of run-away seismic rupture on otherwise unstable faults, even
 642 for sustained increases of fault permeability with slip. On ultimately stable
 643 faults, instead, dilatant hardening effect delays the nucleation of finite-sized
 644 seismic event and increases its dynamic run-out distance. We can certainly
 645 say, however, that the findings obtained in this work are valid for an imma-
 646 ture fault whose dilatant compliance is relatively small, i.e. for $\frac{\Delta w_h}{w_h c_f \sigma'_o} \ll 0.1$

647 which would result in an undrained pore-pressure drop of fractions of MPa.

648 In this contribution we also neglect other weakening mechanisms that
 649 can kick in during the onset of dynamic crack propagation, such as thermal
 650 pressurization and flash heating (Rice, 2006).

651 Thermal pressurization of pore-fluid by rapid shear heating of fault gouge
 652 unit has been showed to be a prominent process of fault weakening (Viesca
 653 and Garagash, 2015). This mechanism depends on fluid thermodynamics
 654 properties and drives the shear strength loss at low fluid pressure conditions
 655 (Acosta et al., 2018). Garagash and Germanovich (2012) showed that dy-
 656 namic weakening due to thermal pressurization increases the dynamic run-out
 657 slip distances on ultimately stable faults. However, they also stated that such
 658 an effect is relatively small, due to the fact the most of dynamic weakening is
 659 expected to occur at slip scale $\delta_{w,dyna}$ that is likely larger than δ_w . Without

660 loss of generality, we can claim that the results presented in this contribution
661 in terms of dynamic run-out distances on ultimately stable faults are valid
662 for sufficiently low values δ_w compared to the characteristic dynamic slip-
663 weakening distance $\delta_{w,dyna}$ (which depends on heat properties of both fault
664 gouge unit and injected fluid).

665 Flash heating on fault's asperities, instead, kicks in when the slip velocity
666 exceeds $\sim 0.1\text{m/s}$, a scenario that is very plausible during dynamic crack
667 propagation, which characteristic slip rate at the rupture tip typically ex-
668 ceeds $\sim 10\text{m/s}$ (Garagash, 2011). Laboratory experiments on various types
669 of rocks have shown a drop of friction coefficient of up to one order of mag-
670 nitude due to flash heating on asperities contacts (Di Toro et al., 2011).
671 These considerations suggest that this weakening mechanism would impact
672 the dynamic run-out distances presented in this contribution. However, this
673 is outside the scope of this work and it is left for future investigations.

674 Finally, we adopted a simple linear slip-weakening friction law compared
675 to a more elaborate Rate- and State- friction model (Dieterich, 1979). Al-
676 though a number of studies have demonstrated that the linear slip-weakening
677 friction model is a good approximation of the velocity weakening R&S friction
678 law (Uenishi and Rice, 2003; Rubin and Ampuero, 2005; Garagash, 2021)
679 even at slip instability (Viesca, 2016b,a), it does not allow to investigate
680 scenarios in which fault frictional properties evolve during pressurization,
681 for instance from velocity weakening to velocity strengthening depending on
682 current slip velocity (as observed in laboratory experiments by Cappa et al.
683 (2019)). Indeed, this can only be captured by using R&S friction law with
684 heterogeneous distribution of a and b parameters. Future works with incor-
685 poration of R&S friction model will follow.

686 *4.2. Implications on injection-induced seismicity*

687 The results above may indicate that injection-induced seismicity can be
688 mitigated by controlling operational parameters, in particular the injection
689 volumetric rate.

690 Upon fluid injection into a specific fracture zone, possibly indicating of a
691 well-developed fault zone, pore-pressure perturbation is likely to activate slip
692 on favourably oriented (and critically stressed) fractures within the damage
693 zone. The initial stable slip propagation could be quickly followed by run-
694 away seismic ruptures that in turn could trigger events on other fractures
695 and propagate over the entire fault zone. Our results show that the onset
696 of the contained micro-seismicity, however, may be delayed in time for slow

697 increases of injection rate prior reaching a steady-state phase. If a specific
698 stable principal fault plane is targeted for fluid injection, instead, a larger
699 pore-fluid perturbation is required to activate slip. A ramp-up of the in-
700 jection rate, in this case, strongly affects the slip propagation, in particular
701 when the ramp-up time scale is much larger than the along-fault diffusion
702 time scale. Although counter-intuitive, a fast ramp-up of injection rate on
703 stable fault planes would reduce the possibility of triggering a larger finite-
704 sized seismic event (with respect to the one that could be triggered if fault
705 pressurization occurs at constant injection rate), which can potentially turn
706 into an unabated rupture due to the activation of other dynamic weakening
707 mechanisms (as discussed in Section 4.1).

708 The results are essentially in line with previous laboratory and numerical
709 observations. Indeed, previous results highlighted how the initial state of
710 effective stress (Gischig, 2015; Passelègue et al., 2018) and essentially its re-
711 lationship to the frictional behaviour (Larochelle et al., 2021) control whether
712 a fault slip is confined to the pressurized region or runaways. Similarly, the
713 numerical analysis by Alghannam and Juanes (2020) reports how the pres-
714 surization rate may influence the seismic reactivation of a fault zone. Here
715 we generalize both approaches by showing that the ruptures and its final
716 behaviour (confined or runaway) are linked to both the initial state of stress
717 as well as the initial ramp-up of the injection, which is then closely linked to
718 the pressurization rate. The effects of the controlling operational parameters
719 on risk of induced-seismicity associated with the specific physic-based model
720 presented in this contribution are summarized in Figure 11. It displays the
721 effects of the initial ramp-up of injection rate on the different slip regimes
722 (the same of Figure 2), as function of all the other governing dimensionless
723 parameters. A comparison with Fig. 2 that is valid for the simple injec-
724 tion scenario of constant rate reveals that, for an ultimately stable fault in
725 which $t_c/t_w \gg 1$, the larger is the pressurization rate (i.e. the larger is the
726 ratio $\frac{Q_m t_w}{t_c Q^*}$), the smaller is region (2) and thus the lower is the possibility of
727 triggering a finite-sized seismic event. For slightly unstable faults ($\tau^o \gtrsim \tau_r$),
728 instead, a lower pressurization rate would lead to a larger slip accumulation
729 during the quick quasi-static phase of crack propagation. The friction coeffi-
730 cient, therefore, would drop quicker to its residual value and the probability
731 of triggering a run-away dynamic rupture preceded by a finite-sized event is
732 higher, resulting in a larger region (3). This also implies that the nucleation
733 time of such a run-away rupture increases for lower pressurization rates, as

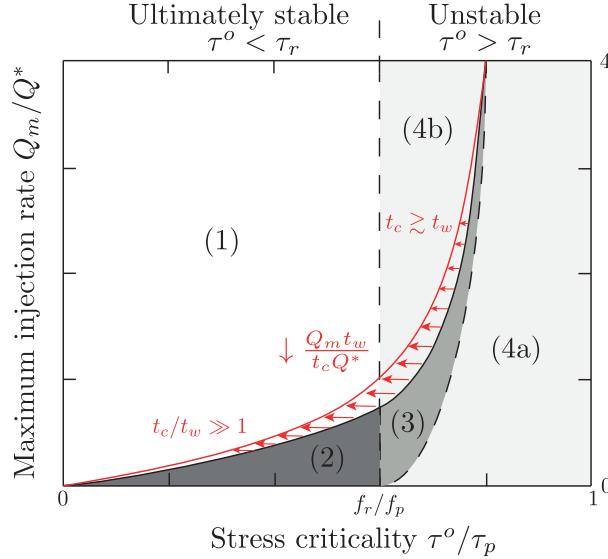


Figure 11: Map of slip regimes for a frictional weakening pressurized fault in which injection rate increases linearly in time, up to reach a maximum plateau after time t_c . This parametric diagram is function of stress criticality τ^o/τ_p , friction weakening ratio f_r/f_p , normalized maximum injection rate Q_m/Q^* and normalized ramp-up time scale t_c/t_w . For a given value of Q_m/Q^* , the larger is the ratio t_c/t_w , i.e. the lower is the pressurization rate $\frac{Q_m t_w}{t_c Q^*}$, the more wide are region (2) and (3). All the slip regimes of this Figure are the same of those reported and described in Figure 2.

734 more and more part of the shear crack has reached the residual shear strength
 735 during its quasi-static stable propagation.

736 5. Conclusions

737 In this study, we have extended the model of Garagash and Germanovich
 738 (2012) to account for an initial ramp-up of injection rate in time before
 739 the maximum plateau and investigated its effect on nucleation and arrest of
 740 dynamic fault slip. Despite the simplicity of the homogeneous model (pla-
 741 nar bidimensional fault, uniform stress conditions and rock properties, linear
 742 weakening of friction, no dilatancy), it allows to get insight into the mech-
 743 anisms that govern the transition between aseismic and seismic slip on a
 744 pressurized planar fault.

745 We have approximated standard injection protocols that consist of an initial
 746 step-wise ramp-up of injection rate in time with a linear increasing function,

747 followed by a maximum plateau after a given time t_c . We have solved nu-
748 merically the coupled hydro-mechanical problem and explored the different
749 slip regimes identified via scaling analysis. Our results show that the initial
750 ramp-up of injection rate affects the slip propagation on ultimately stable
751 faults ($\tau^o < \tau_r$) if and only if the ramp-up time scale t_c is much larger than
752 fluid diffusion time scale t_w . Notably, we have shown that slip stability is gov-
753 erned by the pressurization rate (quantified by the dimensionless parameter
754 $\frac{Q_m t_w}{Q^* t_c}$) and thus by how quickly the injection rate increases before reach-
755 ing time t_c . From our results we can conclude that low pressurization rates
756 applied on ultimately stable faults with frictional weakening properties and
757 $t_c \gg t_w$ promote the nucleation of a finite-sized dynamic event. Moreover,
758 the lower is the pressurization rate, the larger is the dynamic run-out slip
759 distance and hence the larger is the magnitude of the induced seismic event.
760 We have also developed an asymptotic solution in terms of nucleation time
761 that is valid when slip instability follows shortly crack activation and verified
762 it with numerical simulations.

763 When $t_c \ll t_w$, instead, the initial ramp-up of injection rate can be neglected
764 on ultimately stable faults and slip are driven by pressurization at constant
765 volumetric rate. A good agreement with theoretical predictions of Garagash
766 and Germanovich (2012) valid for that particular injection condition has been
767 obtained.

768 Finally, we have demonstrated that the ramp-up of injection rate does affect
769 the quick slip propagation on critically stressed faults prior the nucleation of
770 a run-away dynamic rupture. By using the universal outer asymptotic so-
771 lution of Garagash and Germanovich (2012), we developed a new analytical
772 solution for the nucleation time and verified it with numerical simulations.
773 This solution reveals that the initial ramp-up of injection rate may lead to
774 an earlier or later nucleation of unabated dynamic event, compared to the
775 case of fault pressurization at constant volumetric rate.

776 **Appendix A. Along-fault pore-pressure diffusion: analytical solu-** 777 **tion**

778 Equation (5) or equivalently equation (9) is a parabolic, linear second-
779 order partial differential equation with constant coefficients. It governs the
780 one-dimensional over-pressure diffusion $\bar{p}(x, t) = p(x, t) - p_o$ inside the fault
781 gouge unit with constant hydraulic diffusivity $\alpha [L^2/T]$. Together with the

782 boundary conditions

$$\bar{p}(\pm\infty, t) = 0, \quad -\frac{w_h k_f}{\mu} \frac{\partial \bar{p}}{\partial x} \Big|_{x=0^\pm} = \pm \begin{cases} \frac{Q_m}{2t_c} \cdot t & t \leq t_c \\ \frac{Q_m}{2} & t > t_c \end{cases}, \quad (\text{A.1})$$

783 and the specific initial conditions, it represents a well-posed diffusion problem
784 that can be solved analytically (Carslaw and Jaeger, 1959).

785 Firstly, we solve the diffusion problem for the case of linear ramp-up of
786 injection rate valid for $t \leq t_c$, for which the initial condition reads

$$\bar{p}(x, 0) = 0 \quad (\text{A.2})$$

787 We use Laplace transform in time (with Laplace parameter s) in order
788 to turn equation (9) into an ordinary differential equation (with transformed
789 variables denoted with an hat $\hat{\cdot}$). With the initial condition (A.2), such a
790 subsidiary equation reads

$$\frac{\partial^2 \hat{p}(x, s)}{\partial x^2} - k^2 \hat{p}(x, s) = 0 \quad \text{with} \quad k^2 = \frac{s}{\alpha}, \quad (\text{A.3})$$

791 whose analytical solution is

$$\hat{p}(x, s) = c_1(s) \cdot e^{kx} + c_2(s) \cdot e^{-kx} \quad (\text{A.4})$$

792 $c_1(s)$ and $c_2(s)$ are two constants that can be obtained in closed form
793 using the boundary conditions (A.1) associated only with linear ramp-up of
794 injection rate:

$$c_1(s) = 0, \quad c_2(s) = -\frac{\mu Q_m}{2t_c k_f s^2 w_h k} \quad (\text{A.5})$$

795 By taking the inverse Laplace transform of equation (A.4) with the constants
796 defined in equation (A.5), we get the analytical expression for along-
797 fault over-pressure diffusion reported in Eq. (10) and valid for $t \leq t_c$, i.e.

$$\bar{p}(\xi, t \leq t_c) = \left(\frac{2Q_m \mu \sqrt{\alpha} t^{3/2}}{3k_f \sqrt{\pi} t_c w_h} \right) \cdot \left(e^{-\xi^2} (1 + \xi^2) - \sqrt{\pi} |\xi| \left(\frac{3}{2} + \xi^2 \right) \text{Erfc}(|\xi|) \right), \quad (\text{A.6})$$

798 where $\xi = \frac{x}{\sqrt{4\alpha t}}$ is the along-fault normalized coordinate and Erfc is the
799 complementary error function.

800 We then solve the pore-fluid diffusion problem for time $t > t_c$, taking into
801 account the injection history during the ramp-up phase.
802 In view of problem linearity, the analytical expression for the over-pressure
803 evolution valid for time $t > t_c$ can be obtained by solving the diffusion
804 equation (9) using as initial condition the over-pressure distribution evaluated
805 at time t_c (using Eq. (A.6)), and as boundary conditions the relations (A.1)
806 associated only with constant injection rate. By using the fundamental heat
807 conduction solution valid for an infinite one-dimensional body subjected to
808 an instantaneous point source (also called Green's function) (Carslaw and
809 Jaeger, 1959)

$$G(x - x', t - t') = \frac{1}{\sqrt{4\pi\alpha(t - t')}} e^{-\frac{(x-x')^2}{4\alpha(t-t')}} \quad \text{for } t - t' \geq 0 \quad (\text{A.7})$$

810 and the superimposition principle, we can express the over-pressure evolution
811 valid for time $t > t_c$ via the following analytical formula (Cole et al., 2011):

$$\bar{p}(x, t > t_c) = \int_{-\infty}^{\infty} G(x-x', t-t_c) \cdot \bar{p}\left(\frac{x'}{\sqrt{4\alpha t_c}}, t_c\right) dx' + \frac{\alpha Q_m \mu}{w_k k_f} \int_{t'=t_c}^t G(x, t-t') dt' \quad (\text{A.8})$$

812 Notice that the first term of Eq. (A.8) is the convolution with respect
813 to the variable x of the fundamental solution (A.7) and the over-pressure
814 distribution at time t_c (obtained from (A.6)), while the second term is the
815 contribution due to the constant flux boundary condition at fault centre. If
816 the ramp-up time t_c vanishes (i.e. $t_c = 0$), then fault pressurization would
817 occur at constant injection rate. The first member of equation (A.8) would
818 vanish (due to the initial condition (A.2)) and the analytical solution for the
819 over-pressure evolution along the fault plane upon integration would read as

$$\bar{p}(x, t) = \frac{\alpha Q_m \mu}{w_k k_f} \cdot \frac{t \cdot E_{\frac{3}{2}}\left(\frac{x^2}{4t\alpha}\right)}{2\sqrt{\pi}\sqrt{\alpha t}}, \quad (\text{A.9})$$

820 where $E_n(z)$ is the Exponential integral function, or expressed in a similar
821 form of (A.6) as

$$\bar{p}(\xi, t) = \left(\frac{Q_m \mu \sqrt{\alpha t}}{\sqrt{\pi} k_f w_h}\right) \cdot \left(e^{-\xi^2} - \sqrt{\pi} |\xi| \text{Erfc}(|\xi|)\right) \quad (\text{A.10})$$

822 with $\xi = \frac{x}{\sqrt{4\alpha t}}$.

823 **Appendix B. Scaling analysis**

824 Scaling analysis applied on physics-based models represents a powerful
 825 technique that can help a systematic investigation of all the physical pro-
 826 cesses occurring. For sake of completeness, we report here the dimensionless
 827 solution structure as well as the normalised set of governing equations.

828 Based on previous work of Garagash and Germanovich (2012), we introduce
 829 the following characteristic scales in order to normalize elasticity equation
 830 (1) and shear stress evolution within the crack tips (3):

$$\tau(x, t) = \tau_p \cdot \mathcal{T}(x, t), \quad \delta(x, t) = \delta_w \cdot \Delta(x, t), \quad \bar{p}(x, t) = \sigma'_o \cdot \Pi(x, t) \quad x = a \cdot \mathcal{X} \quad (\text{B.1})$$

831 where δ_w is the slip weakening length-scale (see Figure 1), $\tau_p = f_p \sigma'_o$ is the
 832 peak shear strength at ambient conditions and a is half-length of the shear
 833 crack. The corresponding governing equations upon introduction of (B.1)
 834 read:

$$\mathcal{T}(a\mathcal{X}, t) = \frac{\tau_o}{\tau_p} - \frac{1}{2\pi} \frac{a_w}{a} \int_{-1}^1 \frac{\partial \Delta(a\mathcal{X}, t)}{\partial \zeta} \frac{d\zeta}{(\zeta/a - \mathcal{X})}, \quad (\text{B.2})$$

835

$$\mathcal{T}(a\mathcal{X}, t) = \frac{f(\delta)}{f_p} (1 - \Pi(a\mathcal{X}, t)), \quad (\text{B.3})$$

836 where $a_w = \frac{E_p}{2\tau_p} \delta_w$ is a characteristic length-scale of the slipping patch and
 837 the normalized friction coefficient evolution is defined as

$$\frac{f(\delta)}{f_p} = \begin{cases} 1 - \Delta(a\mathcal{X}, t) & \Delta(a\mathcal{X}, t) \leq 1 - \frac{f_r}{f_p} \\ \frac{f_r}{f_p} & \Delta(a\mathcal{X}, t) > 1 - \frac{f_r}{f_p} \end{cases} \quad (\text{B.4})$$

838 In order to normalize the fluid diffusion problem (5-8), we use the same
 839 characteristic scales introduced in (B.1), with the exception that now we
 840 scale the spatial coordinate x with the length-scale a_w previously obtained.

841 Time t , instead, is normalized with the diffusion time-scale $t_w = \frac{a_w^2}{4\alpha}$ such
 842 that

$$t = t_w \cdot \text{T} \quad (\text{B.5})$$

843 Equations (5-8), therefore, reduce to

$$\frac{\partial \Pi}{\partial \text{T}} - \frac{1}{4} \frac{\partial^2 \Pi}{\partial \mathcal{X}^2} = 0, \quad (\text{B.6})$$

844

$$\frac{\partial \Pi}{\partial \mathcal{X}} \Big|_{\mathcal{X}=0^\pm} = \begin{cases} \pm \frac{Q_m}{Q^*} \frac{t_w}{t_c} \cdot T & T \leq \frac{t_c}{t_w} \\ \pm \frac{Q_m}{Q^*} & T > \frac{t_c}{t_w} \end{cases} \quad (\text{B.7})$$

845 where Q^* is the characteristic scale of maximum injection rate and is defined
846 as

$$Q^* = \frac{2\sigma'_o w_h k_f}{a_w \mu} \quad (\text{B.8})$$

847 Inspecting equations (B.2-B.4) and (B.6-B.7) we can readily observe that
848 the dimensionless solution is function of only four dimensionless parameters:

- 849 • $\frac{\tau^o}{\tau_p} \rightarrow$ fault stress criticality
- 850 • $\frac{f_r}{f_p} \rightarrow$ friction weakening ratio
- 851 • $\frac{Q_m}{Q^*} \rightarrow$ normalized maximum injection rate
- 852 • $\frac{t_c}{t_w} \rightarrow$ normalized ramp-up time scale

853 Appendix C. Small-Scale Yielding asymptotics

854 When half-length of slipping patch a is sufficiently larger than the char-
855 acteristic length-scale a_w , the model non-linearity (i.e. slip weakening of
856 friction coefficient) is localized only over a region that is small compared
857 to geometrical dimensions of the rupture zone. A clear example of such a
858 condition, typically named as small-scale yielding, is depicted in Figure 4-
859 bottom where, after the arrest of dynamic slip, the friction coefficient has not
860 reached its residual value f_r only over a small region near crack tips. The
861 elastic stress-intensity factor, therefore, controls the local deformation field
862 when $a \gg a_w$ (Rice, 1968), and its analytical expression can be obtained in
863 closed form by superimposing the effects of net loading applied on the shear
864 crack (i.e. far-field stress τ^o minus residual shear tractions τ_r) and of fluid
865 pressurization as (Tada et al., 2000)

$$K_{II} = (\tau^o - \tau_r) \sqrt{\pi a} + \Delta K_{II}, \quad (\text{C.1})$$

866 where

$$\Delta K_{II} = f_r \sqrt{\frac{a}{\pi}} \int_{-a}^{+a} \frac{\bar{p}(x, t)}{\sqrt{a^2 - x^2}} dx \quad (\text{C.2})$$

867 By replacing the analytical expression for over-pressure evolution during
 868 the ramp-up phase (10) into (C.2), we find that the increment of stress-
 869 intensity factor due to pressurization at time $t \leq t_c$ is

$$\begin{aligned} \Delta K_{II} = & -\frac{f_r \sqrt{a} \mu Q_m t e^{-\frac{a^2}{8\alpha t}}}{36 \sqrt{\pi} k_f t_c w_h (\alpha t)^{3/2}} \left(4a^3 \sqrt{\alpha t} e^{\frac{a^2}{8\alpha t}} + 36a(\alpha t)^{3/2} e^{\frac{a^2}{8\alpha t}} - \sqrt{\pi} a^4 I_0 \left(\frac{a^2}{8t\alpha} \right) - \right. \\ & - 9\sqrt{\pi} a^2 \alpha t I_0 \left(\frac{a^2}{8t\alpha} \right) - \sqrt{\pi} a^4 I_1 \left(\frac{a^2}{8t\alpha} \right) - 11\sqrt{\pi} a^2 \alpha t I_1 \left(\frac{a^2}{8t\alpha} \right) - \\ & - 3\sqrt{\pi} a^2 \alpha t e^{\frac{a^2}{8\alpha t}} \cosh \left(\frac{a^2}{8|\alpha||t|} \right) I_0 \left(\frac{a^2}{8|t||\alpha|} \right) - \\ & - 24\sqrt{\pi} \alpha^2 t^2 e^{\frac{a^2}{8\alpha t}} \cosh \left(\frac{a^2}{8|\alpha||t|} \right) I_0 \left(\frac{a^2}{8|t||\alpha|} \right) + \\ & + 3\sqrt{\pi} a^2 |\alpha||t| e^{\frac{a^2}{8\alpha t}} \cosh \left(\frac{a^2}{8|\alpha||t|} \right) I_1 \left(\frac{a^2}{8|t||\alpha|} \right) + \\ & + 3\sqrt{\pi} a^2 |\alpha||t| e^{\frac{a^2}{8\alpha t}} \sinh \left(\frac{a^2}{8|\alpha||t|} \right) I_0 \left(\frac{a^2}{8|t||\alpha|} \right) + \\ & + 24\sqrt{\pi} \alpha t |\alpha||t| e^{\frac{a^2}{8\alpha t}} \sinh \left(\frac{a^2}{8|\alpha||t|} \right) I_0 \left(\frac{a^2}{8|t||\alpha|} \right) - \\ & \left. - 3\sqrt{\pi} a^2 \alpha t e^{\frac{a^2}{8\alpha t}} \sinh \left(\frac{a^2}{8|\alpha||t|} \right) I_1 \left(\frac{a^2}{8|t||\alpha|} \right) \right) \quad (\text{C.3}) \end{aligned}$$

870 where $I_n(z)$ is the modified Bessel function of the first kind.

871 If pressurization consists of injection at constant flow rate only, then the
 872 increment of stress-intensity factor would be obtained by replacing equation
 873 (A.10) into (C.2), leading to the following expression

$$\Delta K_{II} = -\frac{e^{-\frac{a^2}{8\alpha t}} \sqrt{a} f_r \mu Q_m}{4\sqrt{\pi} k_f w_h \sqrt{\alpha t}} \left(4a e^{\frac{a^2}{8\alpha t}} \sqrt{\alpha t} - \sqrt{\pi} a^2 I_0 \left(\frac{a^2}{8t\alpha} \right) - \sqrt{\pi} a^2 I_1 \left(\frac{a^2}{8t\alpha} \right) - 4\sqrt{\pi} \alpha t I_0 \left(\frac{a^2}{8t\alpha} \right) \right) \quad (\text{C.4})$$

874 During the slipping patch propagation driven by pore-fluid diffusion, the
 875 energy release rate $G = \frac{K_{II}^2}{E_p}$ must be equal to the fracture energy G_c , whose

876 analytical expression under the assumption of small-scale yielding and con-
 877 stant effective normal stress near crack tips is (Palmer and Rice, 1973)

$$G_c \simeq (f_p - f_r)\sigma'(a)\frac{\delta_r}{2} \quad (\text{C.5})$$

878 The quasi-static propagation criterion

$$G = G_c \quad (\text{C.6})$$

879 therefore provides an asymptotic solution for the slipping patch length a as
 880 function of pressurization time t .

881 Following the previous work of Dempsey et al. (2010), however, Garagash
 882 and Germanovich (2012) have shown that by replacing the slipping patch
 883 length a in (C.1-C.2) by an effective (reduced) length a_{eff} , a more accurate
 884 expression of the stress-intensity factor in the limit of $a \gg a_w$ is obtained.
 885 Such an effective crack length a_{eff} is function of a process zone size d that
 886 Dempsey et al. (2010) determined numerically for the case of cohesive crack
 887 with linear softening traction separation law and propagating under uniform
 888 far-field tractions (which share the same mathematical formulation), i.e.

$$a_{eff} \simeq a - 0.466 \cdot d, \quad (\text{C.7})$$

889 where d is defined as

$$d \simeq 0.466 \cdot \lambda \quad (\text{C.8})$$

890 and $\lambda = \left(\frac{\pi}{2}\right) \cdot \left(\frac{K_{II}}{(\tau_p - \tau_r)}\right)^2$ is a characteristic length-scale.

891 In this contribution, the implicit equation (C.6) incorporates the reduced
 892 crack length (C.7) and it is solved numerically for time t (thus imposing the
 893 variable crack length a) by minimizing the residual function using a random
 894 search method. When the slipping patch is driven essentially by pressuriza-
 895 tion at constant injection rate (for instance in the case of an ultimately stable
 896 fault in the limit when $t_c/t_w \ll 1$), equation (C.6) is solved using the stress-
 897 intensity factor obtained in (C.4). Instead, when the shear crack is driven
 898 up to the nucleation of a dynamic event by the ramp-up of injection rate
 899 (for instance in the case of an ultimately stable fault with $t_c/t_w \gg 1$), then
 900 equation (C.6) is solved using (C.3) (obviously up to the maximum value t_c).
 901 The comparison between the numerical results (black lines) and small-scale
 902 asymptotic solutions (dashed grey lines) in Figures 3 and 6 shows a good
 903 match for slipping patches a larger than $\sim 2a_w$.

904 **Appendix D. Outer and inner asymptotic solutions at instability**
 905 **for an unstable fault**

906 Here we briefly report the outer and inner asymptotic solutions at instabil-
 907 ity for an unstable fault $\tau^o/\tau_r > 0$ developed by Garagash and Germanovich
 908 (2012). The former is universal, which means it is independent of a par-
 909 ticular pore-pressure profile, while the latter does depend on the particular
 910 pore-pressure distribution. Since dynamic instability for an unstable fault
 911 occurs quickly after fluid injection and crack activation (and thus during the
 912 ramp-up of injection rate for values of t_c comparable to or greater than t_w),
 913 the inner asymptotic solution is solved only for the particular distribution
 914 (10).

915 *Appendix D.1. Outer solution*

916 At instability, the slipping patch extent is much larger than the pressur-
 917 ized region, due to the rapid propagation of the shear crack after its activa-
 918 tion. Under this condition, the pore pressure distribution can be replaced by
 919 an equivalent point-force distribution

$$\bar{p}(x, t) \simeq \Delta P(t) \delta_{dirac}(x) \quad (\text{D.1})$$

920 where δ_{dirac} is the delta Dirac function and $\Delta P(t) = \int_{-\infty}^{\infty} \bar{p}(x, t) dx$. This
 921 equivalent distribution, however, approximates well $\bar{p}(x, t)$ only for distances
 922 much larger than fluid front position, i.e. for $x \gg \sqrt{4\alpha t}$ (outside pressuriza-
 923 tion region).

924 Garagash and Germanovich (2012) solved semi-analytically the hydro-mechanical
 925 problem with the equivalent pore-pressure distribution (D.1) and proved that
 926 the normalized slip profile at instability is given by

$$\frac{\delta(x)}{\epsilon \delta_w} = \mathcal{P} \left(\bar{\delta}_1(X) - \frac{a}{a_w} \bar{\delta}_2(X) \right) + \bar{\delta}(X), \quad (\text{D.2})$$

927 where $\mathcal{P} \simeq 0.8369$ is the scaled magnitude of the point force, $\epsilon = 1 - \tau^o/\tau_p$ is
 928 the understress parameter, $X = x/a$ is the normalized fault coordinate, and

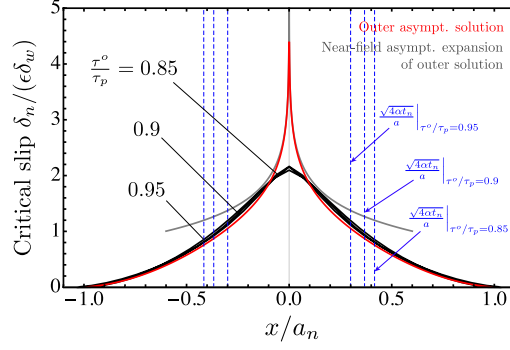


Figure D.12: Corresponding normalized slip distributions $\delta_n/(\epsilon\delta_w)$ (black solid lines) at nucleation time t_n of the numerical simulations reported in Figure 10. The red solid line denotes the outer asymptotic solution valid outside the pressurization regions (whose extents at instability are identified by blue dashed lines), i.e. for $\left|\frac{x}{a_n}\right| \gg \frac{\sqrt{4\alpha t_n}}{a_n}$, while the grey solid line represents the near-field expansion of the outer asymptotic solution that is valid for $x/a_n \ll 1$.

929 $\bar{\delta}_1(X)$, $\bar{\delta}_2(X)$, and $\bar{\delta}(X)$ are three functions that are defined as

$$\begin{aligned}
 \bar{\delta}_1(X) &= \frac{2}{\pi} \cdot \ln \left(\frac{1 + \sqrt{1 - X^2}}{|X|} \right), \\
 \bar{\delta}(X) &\simeq -1.1732 \cdot \sin(\arcsin(X)) - 0.0608 \cdot \sin(3 \cdot \arcsin(X)) + 0.0235 \cdot \sin(5 \cdot \arcsin(X)) \\
 \bar{\delta}_2(X) &= \frac{2}{\pi} \int_{-1}^1 \ln \left| \frac{X - s}{1 - sX + \sqrt{1 - s^2} \sqrt{1 - X^2}} \right| \bar{\delta}_1(s) ds
 \end{aligned} \tag{D.3}$$

930 In Eq. (D.3), $\bar{\delta}(X)$ is a continuous and differentiable function whose
 931 analytical expression has been obtained numerically using Gauss-Chebyshev
 932 polynomial quadrature (with truncation at third term - see (Garagash and
 933 Germanovich, 2012) for more details), while $\bar{\delta}_2(X)$ is the inverse of the
 934 Cauchy integral in terms of $\bar{\delta}(X)$ (solved numerically in this contribution).
 935 Equation (D.2) represents a good approximation of slip distribution only
 936 for $\left|\frac{x}{a}\right| \gg \frac{\sqrt{4\alpha t}}{a}$. A near-field asymptotic expansion of (D.2), however, is
 937 proved to be a good approximation of the critical distribution of normalized
 938 slip $\delta/(\epsilon\delta_w)$ for $X \ll 1$ and $a = a_n$ (Garagash and Germanovich, 2012). Such

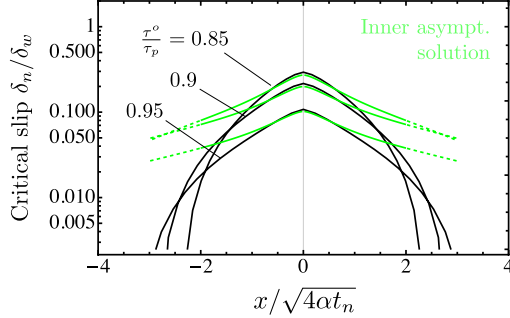


Figure D.13: Linear-log plot showing the corresponding normalized slip distributions δ_n/δ_w (black solid lines) at nucleation time t_n of the numerical simulations reported in Figure 10. The green solid lines denote the inner asymptotic solutions valid within the pressurization regions, i.e. for $\left| \frac{x}{\sqrt{4\alpha t_n}} \right| \lesssim 1$.

939 an expansion is given as

$$\frac{\delta(x)}{\epsilon\delta_w} = -\frac{2}{\pi}\mathcal{P}\ln\frac{|X|}{2} - \frac{a}{a_w}\mathcal{P}\bar{\delta}_2(0) + \bar{\delta}(0) + \mathcal{O}(X), \quad (\text{D.4})$$

940 where the ratio a_n/a_w is defined in (18).

941

942 In order to check the accuracy of these theoretical predictions, we display
 943 in Figure D.12 the corresponding slip profiles at nucleation time t_n of the
 944 numerical simulations reported in Fig. 10. As we can observe, the scaled slip
 945 distributions $\delta_n/(\epsilon\delta_w)$, which collapse into nearly one black line due to the
 946 scaling adopted, match the outer asymptotic solution (D.2) (denoted by a
 947 red solid line) for $|x/a_n| \gg \sqrt{4\alpha t_n}/a_n$, i.e. outside the pressurization regions
 948 whose extents are identified by blue dashed lines (for a given value of stress
 949 criticality τ^o/τ_p). For $x/a_n \ll 1$, instead, the scaled slip distributions tend to
 950 converge to the near-field asymptotic expansion of the outer solution (D.4)
 951 (as expected).

952 *Appendix D.2. Inner solution*

953 Within the small pressurization region, i.e. for $x \lesssim \sqrt{4\alpha t}$, the normalized
 954 pore-pressure distribution is approximately $\bar{p}(x)/\sigma'_o \sim 1$, which means that
 955 the stress perturbation can be written as (Garagash and Germanovich, 2012)

$$\frac{\tau(x) - \tau^o}{\tau_p} \simeq -\frac{\bar{p}(x)}{\sigma'_o} + \mathcal{O}(\epsilon) \quad (\text{D.5})$$

956 for $\epsilon \ll 1$. Using this condition, Garagash and Germanovich (2012) showed
 957 that the normalized slip distribution at instability δ_n/δ_w for a particular
 958 pore-pressure profile $\Psi(s)$ is given as

$$\frac{\delta_n(\xi)}{\delta_w} = \frac{\delta_n(0)}{\delta_w} - \frac{2}{\pi} \epsilon \mathcal{P} \frac{\int_{-\infty}^{\infty} \Psi(s) \ln |1 - \xi/s| ds}{\int_{-\infty}^{\infty} \Psi(s) ds} \quad (\text{D.6})$$

959 where $\delta_n(0)$ is the critical slip at $\xi = \frac{x}{4\alpha t} = 0$, which can be obtained by
 960 matching the outer ($x \gg \sqrt{4\alpha t}$) and inner ($x \lesssim \sqrt{4\alpha t}$) asymptotic solutions
 961 at intermediate distances (Garagash and Germanovich, 2012)

$$\frac{\delta_n(0)}{\delta_w} = 0.533\epsilon \left(-\ln \frac{\hat{\epsilon}}{C} + 1.003 \right), \quad (\text{D.7})$$

962 where $\hat{\epsilon} = (\tau_p - \tau^o)/f_p \Delta \bar{p}(t_n)$ and the constant C is defined as

$$C = \left(\int_{-\infty}^{\infty} \Psi(s) ds \right) \cdot \text{Exp} \left(-\frac{\int_{-\infty}^{\infty} \Psi(s) \ln |s| ds}{\int_{-\infty}^{\infty} \Psi(s) ds} \right) \quad (\text{D.8})$$

963 Unlike the outer asymptotic solution, the inner solution does depend
 964 on the particular pore-pressure profile $\Psi(s)$. By replacing the instanta-
 965 neous spatial distribution (10) associated with ramp-up of injection rate,
 966 we obtain $C = 3.75576$, allowing us to calculate the normalized slip dis-
 967 tribution using (D.7) and (D.6) (with numerical evaluation of the integral
 968 $\int_{-\infty}^{\infty} \Psi(s) \ln |1 - \xi/s| ds$). For the same injection condition, friction weaken-
 969 ing ratio and initial loading conditions of the simulations reported in Figure
 970 10, we show in Figure D.13 the corresponding normalized slip distributions
 971 δ_n/δ_w at nucleation time t_n on a linear-log plot. We can observe that, for each
 972 value of (large) stress criticality τ^o/τ_p , the corresponding inner asymptotic so-
 973 lution matches well the numerical results (back solid lines) for $|x|/\sqrt{4\alpha t_n} \lesssim 1$
 974 (as expected).

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981 The analytical formulae and numerical methods described in the main text
982 and appendices are sufficient to reproduce all the results presented in the
983 paper.

984 **Authors' contributions**

985 F.C. contributed to the Conceptualization, Methodology, Software, Vali-
986 dation, Formal analysis, Investigation, Visualization, Writing original draft.
987 A.P.R. contributed to Editing, Supervision.

988 **Conflict of interest**

989 The authors declare that they have no conflict of interest.

990 **References**

- 991 Acosta, M., Passelègue, F.X., Schubnel, A., Violay, M., 2018. Dynamic
992 weakening during earthquakes controlled by fluid thermodynamics. *Nature*
993 *communications* 9.
- 994 Alghannam, M., Juanes, R., 2020. Understanding rate effects in injection-
995 induced earthquakes. *Nature communications* 11.
- 996 Almakari, M., Dublanchet, P., Hervé, E., Pellet, F., 2019. Effect of the
997 injection scenario on the rate and magnitude content of injection-induced
998 seismicity: Case of a heterogeneous fault. *Journal of Geophysical Research:*
999 *Solid Earth* 124, 8426–8448.
- 1000 Amann, F., Gischig, V., Evans, K., Doetsch, J., Jalali, R., Valley, B.,
1001 Krietsch, H., Dutler, N., Villiger, L., Brixel, B., Klepikova, M., Kittilä,
1002 A., Madonna, C., Wiemer, S., Saar, M., Loew, S., Driesner, T., Mau-
1003 rer, H., Giardini, D., 2018. The seismo-hydromechanical behavior dur-
1004 ing deep geothermal reservoir stimulations: Open questions tackled in a
1005 decameter-scale in situ stimulation experiment. *Solid Earth* 9, 115–137.
1006 doi:10.5194/se-9-115-2018.
- 1007 Barbour, A., Norbeck, J.H., Rubinstein, J.L., 2017. The effects of varying
1008 injection rates in Osage County, Oklahoma, on the 2016 Mw 5.8 Pawnee
1009 earthquake. *Seismological Research Letters* 88, 1040–1053.

- 1010 Batchelor, A.S., 1985. Hot dry rock reservoir stimulation in the UK an ex-
1011 tended summary, in: International seminar on the results of EC geothermal
1012 energy research. 3, European Commission, München. pp. 681–711.
- 1013 Brace, W., Kohlstedt, D., 1980. Limits on lithospheric stress imposed by
1014 laboratory experiments. *Journal of Geophysical Research: Solid Earth* 85,
1015 6248–6252.
- 1016 Cappa, F., Scuderi, M.M., Collettini, C., Guglielmi, Y., Avouac, J.P., 2019.
1017 Stabilization of fault slip by fluid injection in the laboratory and in situ.
1018 *Science advances* 5.
- 1019 Carslaw, H.S., Jaeger, J.C., 1959. *Conduction of heat in solids*. Oxford Univ
1020 Press.
- 1021 Chang, K.W., Yoon, H., Martinez, M.J., 2018. Seismicity rate surge on
1022 faults after shut-in: Poroelastic response to fluid injection. *Bulletin of the*
1023 *Seismological Society of America* 108, 1889–1904.
- 1024 Chen, K.H., Bürgmann, R., 2017. Creeping faults: Good news, bad news?
1025 *Rev. Geoph.* 55, 282–286.
- 1026 Ciardo, F., Lecampion, B., 2019. Effect of dilatancy on the transition from
1027 aseismic to seismic slip due to fluid injection in a fault. *Journal of Geo-*
1028 *physical Research: Solid Earth* 124, 3724–3743.
- 1029 Ciardo, F., Lecampion, B., Fayard, F., Chaillat, S., 2020. A fast boundary
1030 element based solver for localized inelastic deformations. *Int. J. Numer.*
1031 *Meth. Engng.* , 1–23.
- 1032 Cole, K.D., Haji-Sheikh, A., Beck, J.V., Litkouhi, B., 2011. *Heat conduction*
1033 *using Green’s functions*. Taylor & Francis.
- 1034 Deichmann, N., Giardini, D., 2009. Earthquakes Induced by the Stimulation
1035 of an Enhanced Geothermal System below Basel (Switzerland). *Seismo-*
1036 *logical Research Letters* 80, 784–798.
- 1037 Dempsey, D., Riffault, J., 2019. Response of Induced Seismicity to Injec-
1038 tion Rate Reduction: Models of Delay, Decay, Quiescence, Recovery, and
1039 Oklahoma. *Water Resources Res.* 55, 656–681.

- 1040 Dempsey, J.P., Tan, L., Wang, S., 2010. An isolated cohesive crack in tension.
1041 Continuum Mechanics and Thermodynamics 22, 617–634.
- 1042 Di Toro, G., Han, R., Hirose, T., De Paola, N., Nielsen, S., Mizoguchi,
1043 K., Ferri, F., Cocco, M., Shimamoto, T., 2011. Fault lubrication during
1044 earthquakes. Nature 471.
- 1045 Diehl, T., Clinton, J., Kraft, T., Husen, S., Plenkers, K., Guilhelm, A., Behr,
1046 Y., Cauzzi, C., Kästli, P., Haslinger, F., Fäh, D., Michel, C., Wiemer, S.,
1047 2014. Earthquakes in Switzerland and surrounding regions during 2013.
1048 Swiss Geological Society 107, 359–357.
- 1049 Diehl, T., Kraft, T., Kissling, E., Wiemer, S., 2017. The induced earthquake
1050 sequence related to the St. Gallen deep geothermal project (Switzerland):
1051 Fault reactivation and fluid interactions imaged by microseismicity. Jour-
1052 nal of Geophysical Research: Solid Earth 122, 7272–7290.
- 1053 Dieterich, J.H., 1979. Modeling of rock friction: 1. experimental results and
1054 constitutive equations. Journal of Geophysical Research: Solid Earth 84,
1055 2161–2168.
- 1056 Edwards, B., Kraft, T., Cauzzi, C., Kästli, P., Wiemer, S., 2015. Seis-
1057 mic monitoring and analysis of deep geothermal projects in St Gallen and
1058 Basel, Switzerland. Geophys. J. Int. 201, 1022–1039.
- 1059 Ellsworth, W.L., 2013. Injection-induced earthquakes. Science 341.
- 1060 French, M., Zhu, W., Banker, J., 2016. Fault slip controlled by stress path
1061 and fluid pressurization rate. Geophys. Res. Letters 43, 4330–4339.
- 1062 Garagash, D., 2011. On fracture energy of flash heating, in: AGU Fall
1063 Meeting Abstracts, pp. T13A–2360.
- 1064 Garagash, D.I., 2021. Fracture mechanics of rate-and-state faults
1065 and fluid injection induced slip. Phil. Trans. R. Soc. A 379.
1066 <https://doi.org/10.1098/rsta.2020.0129>.
- 1067 Garagash, D.I., Germanovich, L.N., 2012. Nucleation and arrest of dynamic
1068 slip on a pressurized fault. Journal of Geophysical Research 117.

- 1069 Garagash, D.I., Rudnicki, J.W., 2003. Shear heating of a fluid-saturated
1070 slip-weakening dilatant fault zone, 1, limiting regimes. *J. Geophys. Res.*
1071 108.
- 1072 Giardini, D., 2009. Geothermal quake risks must be faced. *Nature* 462,
1073 848–849.
- 1074 Gischig, V., 2015. Rupture propagation behavior and the largest possible
1075 earthquake induced by fluid injection into deep reservoirs. *Geophys. Res.*
1076 *Lett.* 42, 7420–7428.
- 1077 Goertz-Allmann, B.P., Goertz, A., Wiemer, S., 2011. Stress drop variations
1078 of induced earthquakes at the Basel geothermal site. *Geophysical Research*
1079 *Letters* 38.
- 1080 Grawinkel, A., Stockhert, B., 1997. Hydrostatic pore fluid pressure to 9 km
1081 depth - Fluid inclusion evidence from the KTB deep drill hole. *Geophys.*
1082 *Res. Letters* 24, 3273–3276.
- 1083 Grigoli, F., Cesca, S., Rinaldi, A.P., Manconi, A., López-Comino, J.A., Clin-
1084 ton, J.F., Westaway, R., Cauzzi, C., Dahm, T., Wiemer, S., 2018. The
1085 November 2017 Mw 5.5 Pohang earthquake: A possible case of induced
1086 seismicity in South Korea. *Science* 360, 1003–1006.
- 1087 Häring, M.O., Schanz, U., Ladner, F., Dyer, B.C., 2008. Characterisation of
1088 the Basel 1 enhanced geothermal system. *Geothermics* 37, 469–495.
- 1089 Hofmann, H., Zimmerman, G., Zang, A., Min, K.B., 2018. Cyclic soft stim-
1090 ulation (CSS): a new fluid injection protocol and traffic light system to
1091 mitigate seismic risks of hydraulic stimulation treatments. *Geothermal*
1092 *Energy* 6.
- 1093 Holland, A.A., 2013. Earthquakes triggered by hydraulic fracturing in south-
1094 central Oklahoma. *Bulletin of the Seismological Society of America* 103,
1095 1784–1792.
- 1096 Horton, S., 2012. Disposal of hydrofracking waste fluid by injection into
1097 subsurface aquifers triggers earthquake swarm in central Arkansas with
1098 potential for damaging earthquake. *Seismological Research Letters* 83,
1099 250–260.

- 1100 Ito, T., Zoback, M.D., 2000. Fracture permeability and in situ stress to
1101 7 km depth in the KTB Scientific Drillhole. *Geophys. Res. Letters* 27,
1102 1045–1048.
- 1103 Keranen, K.M., Savage, H., Abers, G., Cochran, E.S., 2013. Potentially in-
1104 duced earthquakes in Oklahoma, USA: Links between wastewater injection
1105 and the 2011 Mw 5.7 earthquake sequence. *Geology* 41, 699–702.
- 1106 Keranen, K.M., Weingarten, M., Abers, G., Bekins, B., Ge, S., 2014. Sharp
1107 increase in central Oklahoma seismicity since 2008 induced by massive
1108 wastewater injection. *Science* 345, 448–451.
- 1109 Kim, K.H., Ree, J.H., Kim, Y., Kim, S., Kang, S.Y., Seo, W., 2018. Assess-
1110 ing whether the 2017 Mw 5.4 Pohang earthquake in South Korea was an
1111 induced event. *Science* , eaat6081.
- 1112 Kim, W.Y., 2013. Induced seismicity associated with fluid injection into a
1113 deep well in Youngstown, Ohio. *Journal of Geophysical Research: Solid*
1114 *Earth* 118, 3506–3518.
- 1115 Lagenbruch, C., Zoback, M.D., 2016. How will induced seismicity in Okla-
1116 homa respond to decreased saltwater injection rates? *Science advances*
1117 2.
- 1118 Larochele, S., Lapusta, N., Ampuero, J.P., Cappa, F., 2021. Constraining
1119 fault friction and stability with fluid-injection field experiments. *Geophys.*
1120 *Res. Lett.* 48.
- 1121 Levandowski, W., Weingarten, M., Ill, R.W., 2018. Geomechanical sensitivi-
1122 ties of injection-induced earthquakes. *Geophys. Res. Letters* 45, 8958–8965.
- 1123 Lockner, D.A., Byerlee, J.D., 1994. Dilatancy in hydraulically isolated faults
1124 and the suppression of instability. *Geophysical Research Letters* 21, 2353–
1125 2356.
- 1126 Noël, C., Passelègue, F.X., Giorgetti, C., Violay, M., 2019. Fault reactivation
1127 during fluid pressure oscillations: transition from stable to unstable slip.
1128 *Journal of Geophysical Research: Solid Earth* 124.
- 1129 Palmer, A.C., Rice, J.R., 1973. The growth of slip surfaces in the progressive
1130 failure of over-consolidated clay, in: *Proceedings of the Royal Society of*

- 1131 London A: Mathematical, Physical and Engineering Sciences, The Royal
1132 Society. pp. 527–548.
- 1133 Passelègue, F.X., Almakari, M., Dublanchet, P., Barras, F., Fortin, J., Violay,
1134 M., 2020. Initial effective stress controls the nature of earthquakes. *Nature*
1135 *communications* 11.
- 1136 Passelègue, F.X., Brantut, N., Mitchell, T.M., 2018. Fault reactivation by
1137 fluid injection: Controls from stress state and injection rate. *Geophys.*
1138 *Res. Lett.* 54, 12837–12846.
- 1139 Qin, Y., Chen, X., Walter, J.I., Haffener, J., Trugman, D.T., Carpenter,
1140 B.M., Weingarten, M., Kolawole, F., 2019. Deciphering the stress state
1141 of seismogenic faults in Oklahoma and Southern Kansas based on an im-
1142 proved stress map. *Journal of Geophysical Research: Solid Earth* 124.
- 1143 Quarteroni, A., Sacco, R., Saleri, F., 2000. *Numerical Mathematics. Texts*
1144 *in applied mathematics*, Springer.
- 1145 Rice, J.R., 1968. Mathematical analysis in the mechanics of fracture, in: H.
1146 Liebowitz (Ed.), *Fracture: An Advanced Treatise*. volume 2. chapter 3,
1147 pp. 191–311.
- 1148 Rice, J.R., 1980. The mechanics of earthquake rupture. *Physics of the Earth’s*
1149 *Interior* , 555–649.
- 1150 Rice, J.R., 1992. Fault Stress States, Pore Pressure Distributions, and the
1151 Weakness of the San Andreas Fault. *International Geophysics* 51, 475–503.
- 1152 Rice, J.R., 2006. Heating and weakening on faults during earthquake slip. *J.*
1153 *Geophys. Res* 111.
- 1154 Rubin, A., Ampuero, J.P., 2005. Earthquake nucleation on (aging) rate and
1155 state faults. *Journal of Geophysical Research: Solid Earth* 110.
- 1156 Rudnicki, J.W., 1979. The stabilization of slip on a narrow weakening fault
1157 zone by coupled deformation-pore fluid diffusion. *Bull. Seism. Soc. Am.*
1158 69, 1011–1026.
- 1159 Rudnicki, J.W., Zhan, Y., 2020. Effect of pressure rate on rate and state
1160 frictional slip. *Geophys. Res. Letters* 47.

- 1161 Samuelson, J., Elsworth, D., Marone, C., 2009. Shear-induced dilatancy
1162 of fluid-saturated faults: Experiment and theory. *Journal of Geophysical*
1163 *Research: Solid Earth* 114.
- 1164 Segall, P., Rice, J.R., 1995. Dilatancy, compaction, and slip instability of a
1165 fluid-infiltrated fault. *Journal of Geophysical Research: Solid Earth* 100,
1166 22155–22171.
- 1167 Segall, P., Rubin, A., Rice, J.R., Schmitt, S., 2008. Dilatancy stabilization
1168 vs thermal pressurization as a mechanism for controlling slow vs fast slip.
1169 AGU Fall Meeting Abstracts .
- 1170 Simpson, D.W., 1986. Triggered earthquakes. *Annu. Rev. Earth Planet. Sci.*
1171 14, 21–42.
- 1172 Sumy, D.F., Cochran, E.S., Keranen, K.M., Wei, M., Abers, G., 2014. Ob-
1173 servations of static Coulomb stress triggering of the November 2011 M5.7
1174 Oklahoma earthquake sequence. *Journal of Geophysical Research: Solid*
1175 *Earth* 119, 1904–1923.
- 1176 Suppe, J., 2014. Fluid overpressures and strength of the sedimentary upper
1177 crust. *Journal of Structural Geology* 69, 481–492.
- 1178 Tada, H., C., P.P., Irwin, G.R., 2000. *The Stress Analysis of Cracks Hand-*
1179 *book*. 3rd ed.
- 1180 Uenishi, K., Rice, J.R., 2003. Universal nucleation length for slip-weakening
1181 rupture instability under nonuniform fault loading. *Journal of Geophysical*
1182 *Research: Solid Earth* 108.
- 1183 Viesca, R.C., 2016a. Self-similar slip instability on interfaces with rate-and
1184 state-dependent friction. *Proc. Roy. Soc. Lond. A* 472, 20160254.
- 1185 Viesca, R.C., 2016b. Stable and unstable development of an interfacial sliding
1186 instability. *Physical Review E* 93, 060202.
- 1187 Viesca, R.C., Garagash, D.I., 2015. Ubiquitous weakening of faults due to
1188 thermal pressurization. *Nature Geoscience* 8, 875.
- 1189 Villiger, L., Gischig, V.S., Doetsch, J., Krietsch, H., Dutler, N.O., Jalali,
1190 M., Valley, B., Selvadurai, P.A., Mignan, A., Plenkers, K., Giardini, D.,

- 1191 Amann, F., Wiemer, S., 2020. Influence of reservoir geology on seismic
1192 response during decameter-scale hydraulic stimulations in crystalline rock.
1193 *Solid Earth* 11, 627–655.
- 1194 Wang, L., Kwiatek, G., Rybacki, E., Bonnelye, A., Bohnhoff, M., Dresen,
1195 G., 2020. Laboratory study on fluid-induced fault slip behavior: The role
1196 of fluid pressurization rate. *Geophysical Research Letters* 47.
- 1197 Warpinski, N.R., Teufel, L.W., 1987. Influence of geologic discontinuities
1198 on hydraulic fracture propagation. *Journal of Petroleum Technology* 39,
1199 209–220.
- 1200 Weingarten, M., Ge, S., Godt, J.W., Bekins, B.A., Rubinstein, J.L., 2015.
1201 High-rate injection is associated with the increase in U.S. mid-continent
1202 seismicity. *Science* 348, 1336–1340.
- 1203 Wibberley, C.A.J., 2002. Hydraulic diffusivity of fault gouge zones and im-
1204 plications for thermal pressurization during seismic slip. *Earth Planets*
1205 *Space* 54, 1153–1171.
- 1206 Wibberley, C.A.J., Yielding, G., Di Toro, G., 2008. Recent advances in
1207 the understanding of fault zone internal structure: A review. *Geological*
1208 *Society Special Publication* 299, 5–33.
- 1209 Wong, T.F., 1986. On the normal stress dependence of the shear fracture
1210 energy. *Earthquake Source Mechanics, Geophysical Monograph Series* 37.
- 1211 Yeo, I.W., Brown, M.R.M., Ge, S., Lee, K.K., 2020. Causal mechanism
1212 of injection-induced earthquakes through the Mw 5.5 Pohang earthquake
1213 case study. *Nature communications* 11.
- 1214 Zhang, X., Jeffrey, R.G., Llanos, E.M., 2005. On plane-strain fluid-driven
1215 shear fracture propagation in elastic solids. *Geophys. J. Int.* 163, 419–430.
- 1216 Zoback, M.D., Zoback, M.L., Mount, V.S., Suppe, J., Eaton, J.P., Healy,
1217 J.H., Oppenheimer, D., Reasenber, P., Jones, L., Raleigh, C.B., Wong,
1218 I.G., Scotti, O., Wentworth, C., 1987. New evidence on the state of stress
1219 of the San Andreas fault system. *Science* 238, 1105–1111.