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# **TEMPERATURE, SALINITY AND OXYGEN CLIMATOLOGY OF AFRICAN CONTINENTAL SHELF WATER, SOUTH OF THE EQUATOR, AND CHANGES SINCE 1945**

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## **ABSTRACT**

The first temperature, salinity and oxygen climatologies for waters of the continuous southern African continental shelf is presented. It is based on oceanographic data collected since 1945, sub-sampled at depths of 5, 50 and 100 m on a mixed-spatial grid with 0.25° to 0.5° resolution. The climatologies capture spatial heterogeneities and seasonal variability in key ocean variables for the southern African shelf in unique detail. The results correspond relatively well with biogeographic boundaries informed by classification schemes grounded in taxonomy, but questions the value of the Large Marine Ecosystem approach. Analysis of decadal trends demonstrates the inherent complexity and spatial heterogeneity associated with environmental variability, and suggest the possibility that decadal periodicities are in the process of being disrupted by a longer-term trend. The overall pattern is that southern African West and South coast shelf waters are becoming warmer, except for some upwelling areas, where cooling is evident. Benguela and Agulhas Bank shelf water are also becoming more oxygen depleted.

## 1 **1. INTRODUCTION**

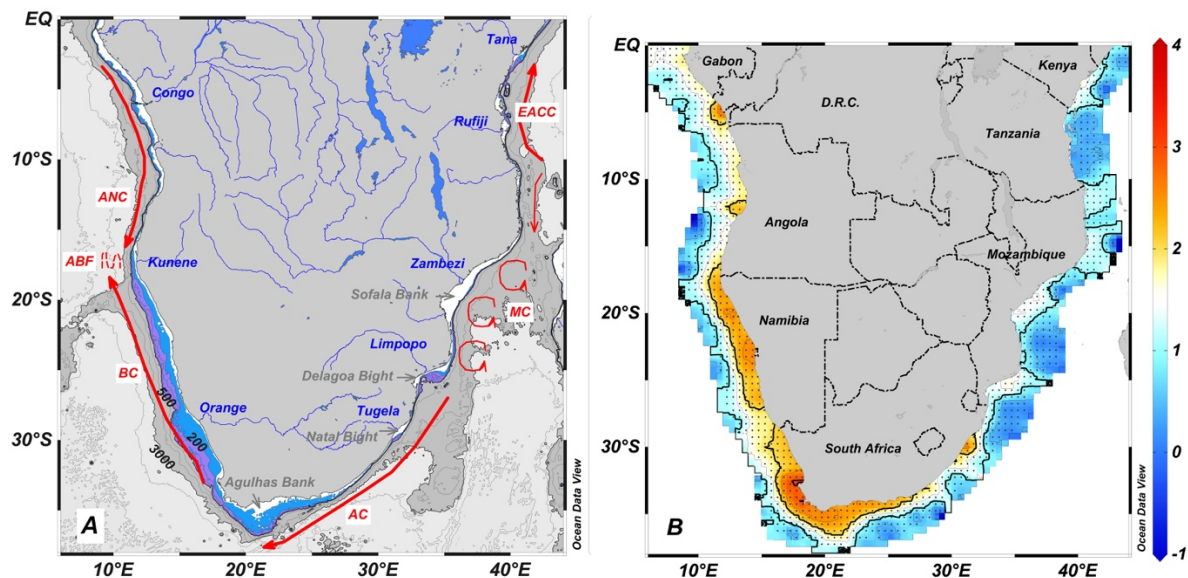
2 Continental shelves account for only about 8% of the world's marine areas, but are host to more marine  
3 biodiversity, productivity and human activities than offshore waters, and provide ecosystem services  
4 comparable in magnitude to that of all terrestrial habitats (UNEP, 2006; Costanza et al., 1997). The  
5 African continental shelf is host to natural resources of particular and critical socio-economic value for  
6 coastal communities (Kainge et al., 2020). Understanding the significance of observed changes in  
7 ocean variables such as temperature and wind (Roualt et al., 2010; Leduc et al., 2010; Santos et al.,  
8 2012; Beal et al., 2016; Vizy et al., 2018; Malan et al., 2019), and ecosystem changes such as species  
9 range shifts and variability in species abundance (Cockcroft et al., 2008; Roy et al., 2007; Coetzee et  
10 al., 2008; Yemane et al., 2014; Blamey et al., 2015; Jarre et al., 2015a; Van der Linge et al., 2016;  
11 Kainge et al., 2017; Van der Linge and Hampton, 2018), have transdisciplinary significance (Potts et  
12 al., 2014; Augustyn et al., 2017; Hobday and Pecl, 2014; Kainge et al., 2020). Construction of  
13 climatological maps of key ocean variables, based on the calculation of mean fields from historical data  
14 sets, are valuable reference tools in this regard.

15 Only a limited number of climatologies have been constructed specifically for shelf areas, e.g. for  
16 the eastern seaboard of North America (Atkinson et al., 1983; Blanton et al., 2003; Bisagni, 2016;  
17 Richaud et al., 2016) and the Ross Sea off Antarctica (Russo et al., 2011). Most climatologies of key  
18 ocean variables exclude data from "land grid-boxes", i.e. coastal areas and portions of the inner  
19 continental shelf (Levitus, 1982; Levitus and Boyer, 1994; Ridgway et al. 2002; Boyer et al. 2005; Boyer  
20 et al. 2018). Large parts of the African shelf is relatively data poor (Boyer et al., 2018), and irregular  
21 and infrequent sampling presents unique challenges to the production of regional scale climatological  
22 maps. This study presents the first such climatological maps for waters of the southern African  
23 continental shelf, from the equator off the west coast of Gabon, to the equator and the east coast of  
24 Somalia.

25 Our knowledge of the morphology of the African continental shelf is much less advanced than for  
26 most other continents (Chiocci and Chivas, 2014). The shelves of the West, South and East African  
27 continental margins comprise passive continental margins characterised by thick sediment  
28 accumulations where fluvial supply is high, such as off the mouths of the Congo, Orange, Tugela,  
29 Limpopo, Save and Zambezi rivers (Scrutton, 1982). The coastline is dominated by lowland coasts  
30 with long unbroken sandy beaches (Orme, 1982); rugged mountainous coasts are rare, except along  
31 the most southern margin. The southern Africa continental shelf varies considerably in width, as defined  
32 by the offshore distance of the 200 m depth isobath (Fig 1A). In the south, the Agulhas Bank has a  
33 maximum offshore extent of about 270 km (Dingle et al., 1978). The shelf areas off Namibia and South  
34 Africa's west coast are also notably wider, compared to those off Angola (Bremner, 1981) and most of  
35 the East Coast (Schumann, 1998). The shelf regions off eastern South Africa and off northern  
36 Mozambique are particularly narrow, often less than 10 km wide, with steep offshore slopes  
37 (Schumann, 1998).

38 The southern Africa shelf is influenced by four surface ocean current systems, the Angola,  
39 Benguela, Agulhas and East African Coastal currents (Fig 1A). The warm southward flowing Angola  
40 Current derives primarily from the South Equatorial Under Current (SEUC) and the Guinea-Congo  
41 Under Current (GCUC) with additional austral winter inflow from the South Equatorial Counter Current  
42 (SECC) (Stramma and Schott, 1999; Mercier et al., 2003). The Angola Current forms the eastern

1 boundary of the Angola Gyre, which is a large scale current field centered near 17°S, 5°E (Gordon and  
 2 Bosley, 1991). Embedded in the gyre is the Angola Dome, centered at 10°S, 9°E (Mazeika, 1967).  
 3 The Angola Current splits into two at the southern boundary of the Angola Gyre at ~ 16 to 17°S: a  
 4 westward continuation, which closes the Angola gyre, and the Benguela Poleward Under Current  
 5 (BPUC), which flows along the upper slope as far south as 27°S off Namibia (Mercier et al, 2003;  
 6 Tchupalanga et al., 2018). At the confluence of the Angola Current and the cold northward flowing  
 7 Benguela Current is a sharp thermal front, the Angola-Benguela Front (ABF, Fig 1A). The ABF migrates  
 8 in an eastward and southward direction in response to the late austral summer (Feb - April) relaxation  
 9 of equatorial easterly winds (Dias, 1983; Stramma and Schott, 1999; Ekau and Verheye, 2005). The  
 10 cool Benguela Current flows northward off the west coast of South Africa and Namibia (Fig 1A), from  
 11 approximately 35°S to just south of the ABF. It is driven by the prevailing South Easterly trade winds  
 12 of the South Atlantic Ocean. Inshore of the Benguela Current, the south easterly winds cause inner  
 13 shelf upwelling of cold nutrient-rich water (Olivar and Shelton, 1993; Shannon, 1995).  
 14



15 Fig 1. A: Bathymetric map of the study area; the coast to 100 m depth interval is filled with white,  
 16 100 to 200 m blue, 200 to 500 m purple (thin solid lines indicate 300 and 400 m depth intervals); 500 to  
 17 200 to 500 m purple (thin solid lines indicate 300 and 400 m depth intervals); 500 to  
 18 3000 m dark grey (thin solid lines indicate depths at 500 m intervals), depths below 3 000 m are light  
 19 grey (with thin solid lines indicating 500 m depth intervals). Also shown are the major continental river  
 20 systems. B: Total number of sampling stations for which oceanographic data is publicly available (as  
 21 WOD and SADCO datasets) for 1900 to 2020, plotted as log(#stations) per 0.5° latitude x 0.5° longitude  
 22 bin. The black dots indicate the center points of the 0.5° x 0.5° degree bins used in this study.  
 23

24  
 25 The warm Agulhas Current dominates the shelf areas off the south and east coasts of South Africa  
 26 (Fig 1A). The current originates near ~ 28°S off Mozambique and transports warm water in a south-  
 27 westerly direction towards its retroflexion area off the tip of Africa (Lutjeharms and van Ballegooyen,  
 28 1988; Lutjeharms and Cooper, 1996). The Agulhas and Benguela currents play a significant role in  
 29 global ocean heat transfer between the Southern and the Northern Hemisphere (Duncombe Rae,  
 30 1991). The ocean area between ~ 10°S and 28°S along the east African coast is occupied by the  
 31 Mozambique Current, a discontinuous current characterised by southward drifting anti-cyclonic eddies  
 32 (Lutjeharms et al., 2012). The East African Coastal Current, or Zanzibar Current, flows northward from

1 ~ 10°S along the east African coast (Swallow et al., 1991). There is close interaction in many cases  
 2 between shelf and slope morphology and current movement, as has been described for south-eastern  
 3 Africa (Lutjeharms 2006).

4 Several biogeographic divisions have been proposed for the southern Africa continental shelf and  
 5 offshore ocean areas. The large marine ecosystem (LME) approach is widely used by regional  
 6 scientists and resource managers (Sherman, 1993), despite its "amorphous" nature (Pauly, 1998). The  
 7 study area hosts the Benguela Current Large Marine Ecosystem (BCLME), the Agulhas Current LME  
 8 (ACLME) and the Somali Coastal Current LME (SCCLME) (Fig 2).

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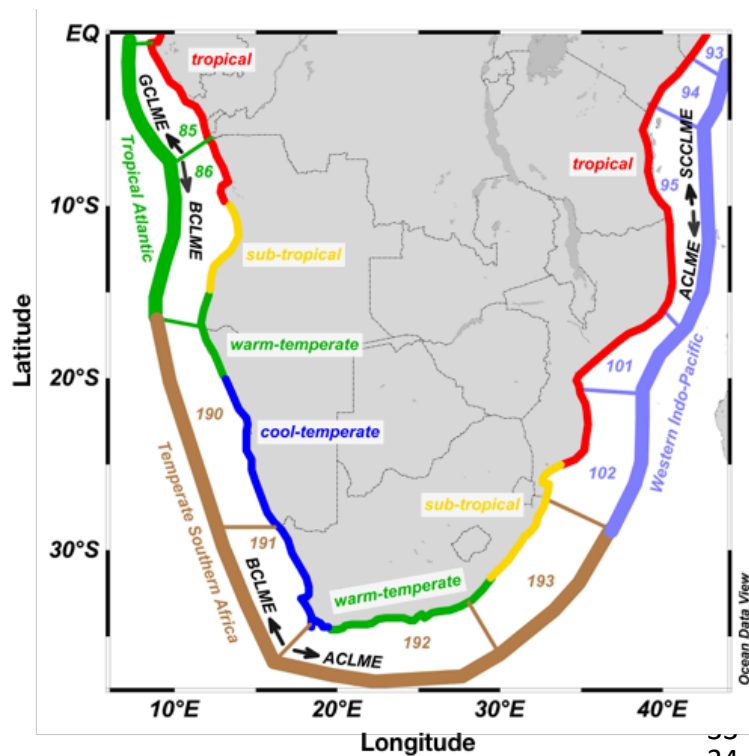


Fig 2. Map of the Tropical Atlantic (thick green line), Temperate Southern Africa (thick brown line) and Western Indo-Pacific (thick purple line) biogeographic Realms of Spalding et al. (2007), subdivided into these *Provinces* (and numbered ecoregion blocks): *Gulf of Guinea* (85 - Gulf of Guinea South, 86 - Angolan), *Benguela* (190 - Namib, 191 - Namaqua), *Agulhas* (192 - Agulhas Bank, 193 - Natal), *Western Indian Ocean* (102 - Delagoa, 101 - Bight of Sofala/Swamp Coast, 95 - East African Coral Coast, 94 - Northern Monsoon Current Coast), *Somali/Arabian* (93 - Central Somali Coast). Large marine ecosystems are indicated in black (LME's; Sherman, 1993): GCLME - Gulf of Guinea Current LME, BCLME - Benguela Current LME, ACLME - Agulhas Current LME, SCCLME - Somali Coastal Current LME. Thin

35 colored lines along the coastline indicate the biogeographical zones defined by Briggs and Bowen  
 36 (2012) and Whitfield (2005), as modified by Potts et al. (2015).  
 37

38 The BCLME is one of the most productive ocean ecosystems in the world in terms of biomass  
 39 production and fishery resources (Spalding et al., 2012). The relatively narrow Northern Benguela shelf,  
 40 off Angola, is characterized by spatially extensive upwelling. The wide Central Benguela (off Namibia)  
 41 and South Benguela (off South Africa) shelf areas, in contrast, host several discrete upwelling cells  
 42 (Shannon, 1995). The hierarchical biogeographic classification system of Spalding et al. (2007) was  
 43 developed specifically and uniquely for the benthic and pelagic biotas of coastal and shelf areas. This  
 44 nested system describes three Realms, five Provinces and eleven biogeographic Ecoregions for the  
 45 southern Africa continental shelf (Fig 2). Other classification systems that rely primarily on taxonomy  
 46 (Briggs and Bowen, 2012; Potts et al. 2015) propose up to 7 biogeographic zones for the African shelf  
 47 and coastline south of the equator (Fig 2).

48

49 Temperature has been proposed to be the key environmental variable correlating with the  
 50 distribution of species and range shifts in shelf areas, as is the case for the open ocean (Spalding et  
 51 al., 2007; Potts et al., 2015; Sunday et al., 2012; Cheung et al., 2013; Morley et al., 2018). It is also

1 increasingly recognized that temperature extremes rather than average values determine population  
2 survival at the warm edge of a species' temperature range (Pinsky et al., 2019). Additionally, abrupt  
3 ecosystem change can be caused by extreme climatic events such as marine heatwaves, that is  
4 prolonged periods of anomalously warm seawater (Shannon et al., 1986; Gammelsrød et al., 1998;  
5 Mills et al., 2012; Cavole et al., 2016; Wernberg et al., 2016; Rouault et al., 2017; Jones et al., 2019;  
6 Smale et al., 2019). Globally, marine heatwaves have increased in frequency and duration since the  
7 early 20th century (Oliver et al., 2018), and global models project this trend to continue through the 21st  
8 century (Oliver et al., 2019). Temperature and dynamic temperature ranges, however, cannot be  
9 considered is isolation of other variables that influence marine species distribution, dissolved oxygen is  
10 also particularly important (Jarre et al., 2015b; Hamukuaya et al., 1998; Mbatha et al., 2019; Woodhead  
11 et al., 1997). This study presents annual and seasonal climatologies for temperature, salinity and  
12 oxygen for the continuous southern African continental shelf at depths of 5, 50 and 100 m. Interdecadal  
13 changes are also evaluated and discussed in the context of documented complex ecosystem changes  
14 that have been observed in the study area.

15

## 16 **2. DATA AND METHODS**

### 17 *2.1. Study area, data sources and distribution*

18 For the purpose of this study, data was extracted for the area 0 to 40°S, 5 to 45°E, from two open  
19 access online sources: WOD18 (World Ocean Database 2018, available at [ncei.noaa.gov](https://ncei.noaa.gov); Boyer et  
20 al., 2018) and SADCO (Southern African Data Centre for Oceanography, presently hosted at  
21 [sadco.ocean.gov.za](https://sadco.ocean.gov.za)). The databases contain both high resolution CTD (i.e. Conductivity-Temperature-  
22 Depth data at 2 or 5 m depth intervals) and low resolution bottle (called OSD or Ocean Station Data in  
23 the WOD) data sets. WOD OSD data was retrieved at standard (as opposed to observed) depths of 0,  
24 5, 50 and 100. The data sets were combined and stations with repeated timestamps removed. The  
25 consolidated WOD-SADCO data set comprises 77 259 individual stations within the 0 to 40°S, 5 to  
26 45°E area. About 85% (65 852) of these stations are located within 3 degrees of the African coastline  
27 (Appendix 1).

28 An outlier filter (Richaud et al., 2016) was applied to this data set as a quality control measure. For  
29 this purpose, all data set values for temperature, salinity and oxygen were grouped into 1° x 1° bins,  
30 and the mean and standard deviation ( $\sigma$ ) calculated for each variable at the surface (0 and 5), 50 and  
31 100 m. All values more than 3 times the standard deviation away from the mean were considered  
32 outliers and removed, with the exception of surface salinity in bins adjacent to river mouths. Only one  
33 iteration of this procedure was carried out, no further data exclusion took place and less than 0.2% of  
34 the total data was removed through this procedure (listed in Appendix 2). The data set was also  
35 screened for spatial and temporal clustering (defined as multiple data points within a 24 h period and 5  
36 km radius) as a precautionary measure. It is important to note that historical data, in particular  
37 biogeochemical data such as oxygen, "have been measured using a variety of manual and automated  
38 analytical techniques" (Boyer et al., 2018) and the details of these are often not contained in metadata.  
39 Comparison of measurement techniques and uncertainties fall outside the scope of the WOD and also  
40 this study.

41 The spatial and temporal coverage of the data set, excluding the above outliers, was then evaluated  
42 by counting the number of OSD and CTD casts in 0.5° x 0.5° bins (Fig 1B, Appendix 1), for a 3 degree

1 distance away from the coastline. This was repeated for each of the variables (Appendix 1). Although  
2 the continental shelf is the focus of this study, it seemed sensible to include data from offshore of the  
3 200 m depth isobath. The  $0.5^\circ \times 0.5^\circ$  bins closest to the coast, i.e. along the inner shelf, were termed  
4  $H_{\text{coast}}$  (or  $H_{0.5^\circ}$ ). Successive bins further offshore ( $H_{1^\circ}$ ,  $H_{1.5^\circ}$ ,  $H_{2^\circ}$ ,  $H_{2.5^\circ}$  and  $H_{3^\circ}$ ) were aligned in an east-  
5 west direction for the West (west of  $18.5^\circ\text{E}$ ) and East (east of  $26.5^\circ\text{E}$ ) coast areas, and in a south-north  
6 direction along the South ( $18.5$  to  $26.5^\circ\text{E}$ ) coast. The  $H_{\text{coast}}$  bins all contain terrestrial land area, to  
7 different degrees. The distribution of data with depth (5, 50 and 100 m), time of day (four 6 hour intervals  
8 from 00:00-05:59 to 18:00-23:50) and time (5 year increments from 1945 to 2019) was established per  
9  $0.5^\circ \times 3^\circ$  bin, using temperature cast data (Appendix 1). In summary, 100% of the total casts contained  
10 temperature data, with equivalent values of 91% for salinity, 54% for oxygen. Additionally, 99% of the  
11 casts contained temperature data at the surface, 49% at 50 m and 37% at 100 m depth.

12

## 13 2.2. Analysis of intra-annual variability and seasonality in $H_{\text{coast}}$ bins

14 Average monthly temperature values were calculated at depths of 5, 50 and 100 m for each  $H_{\text{coast}}$   
15 bin, from  $0$  to  $34.5^\circ\text{S}$  along the West Coast,  $18.5$  to  $26.5^\circ\text{E}$  along the South Coast and  $34.5$  to  $0^\circ\text{S}$  along  
16 the East Coast (Fig 3). Constructing an annual profile consisting of monthly values (with  $n > 3$ ) was  
17 possible for most of the  $H_{\text{coast}}$  bins, with minimal interpolation required, except for the East Coast and  
18 in particular north of  $28^\circ\text{S}$  along the East Coast (Table 1; Fig 3). In the latter instance, monthly data  
19 gaps were filled with the average of four values: that of the preceding and following month of the same  
20 bin, and values for the same month in the two adjacent  $H_{\text{coast}}$  bins. Temperature required the least  
21 amount of gap-filling, and typically up to a factor of two less than salinity and oxygen. In an offshore  
22 direction, i.e.  $H_{\text{coast}}$  towards  $H_{3^\circ}$ , the amount of gap-filling required for the construction of annual profiles  
23 consisting of monthly values is excessive in most areas, even for temperature (Table 1). The method  
24 adopted by Ridgway et al. (2002) and Richaud et al. (2016), that is of assuming an annual profile and  
25 then using it to calculate monthly values for gap-filling purposes, and calculating an annual  
26 climatological average from that, was therefore not adopted in this study. This study, rather, constructed  
27 an annual climatology from four seasonal values. The choice of months to be grouped together to  
28 constitute a season, was based on the intra-annual variability observed in the  $H_{\text{coast}}$  bins, as detailed  
29 below.

30 Intra-annual surface temperature variability in  $H_{\text{coast}}$  bins indicated that the warmest month of the  
31 year, along the entire southern African shelf, is most often February (55%), followed by March (34%)  
32 and January (11%) (Fig 3A, 4A). At the surface, the coldest month is most often August (46 %, blue  
33 line in Fig 3A; Fig 4B), followed by July (23%) and September (20%). At 50 and 100 m depths, however,  
34 the warmest and coldest monthly values were generally observed at different times of the year than at  
35 the surface and the seasonal pattern is more complex and geographically variable than at the surface  
36 (Fig 3, 4). February is the warmest month of the annual cycle for 20 % and 0% of the  $H_{\text{coast}}$  bins at 50  
37 and 100 m, respectively, and August the coldest month for 24% and 22% of the  $H_{\text{coast}}$  bins, at these  
38 two deeper depths. The geographic complexity in intra-annual variability were captured in shelf  
39 segment composites, consisting of the average of all the  $H_{\text{coast}}$  bin monthly values in that segment;  
40 Gabon-Angola or WC:  $0 - 17^\circ\text{S}$ , Central Benguela or WC:  $17 - 28^\circ\text{S}$ , Southern Benguela or WC:  $28 -$   
41  $34^\circ\text{S}$ , Agulhas Bank or SC:  $18.5 - 26.5^\circ\text{E}$  and the southern East Coast or EC:  $25 - 34^\circ\text{S}$  (Fig 5). Based  
42 on the results of this evaluation, months were grouped together for the construction of seasonal average

1 temperature, salinity and oxygen values, as follows: Jan-Feb-Mar (JFM) = summer; Apr-May-Jun (AMJ)  
 2 = autumn; Jul-Aug-Sep (JAS) = winter; and Oct-Nov-Dec (OND) = spring.  
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 6 Fig 3. Comparative average monthly temperature values at depths of 5, 50 and 100 m (A to C) for  $H_{\text{coast}}$   
 7 bins, from 0 to 34.5°S along the West Coast, 18.5 to 26.5°E along the South Coast and 34.5 to 0°S  
 8 along the East Coast. The red lines highlight values for February, most often the warmest month of the  
 9 year in the study area, and the blue lines highlight values for August, most often the coldest month of  
 10 the year. Grey lines indicate the other ten months of the year.

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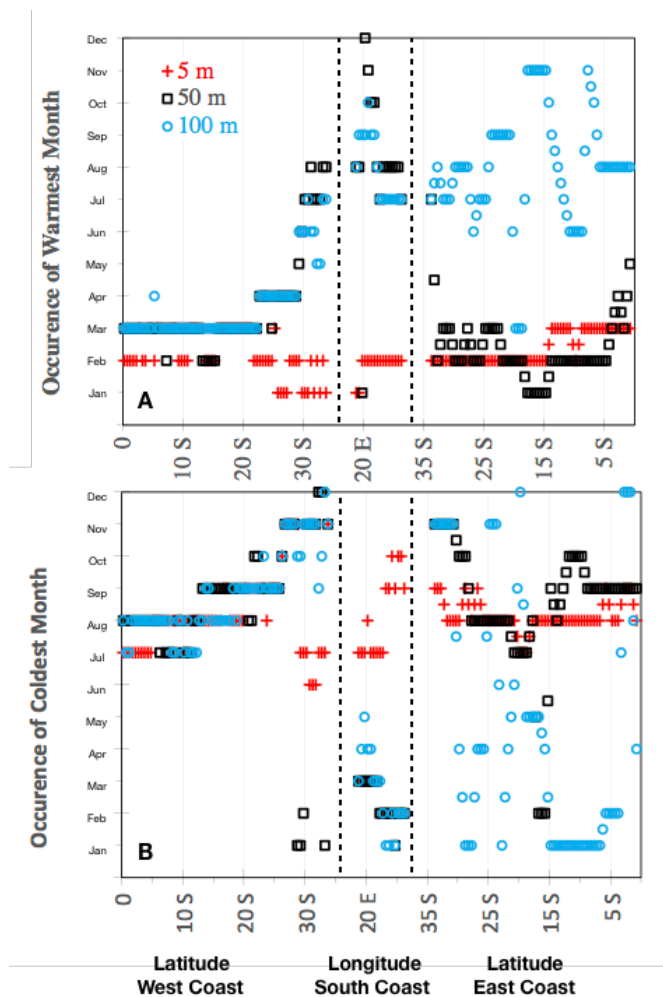


Fig 4. Months in which the highest (A) and lowest (B) monthly average T values are observed in  $H_{\text{coast}}$  bins.

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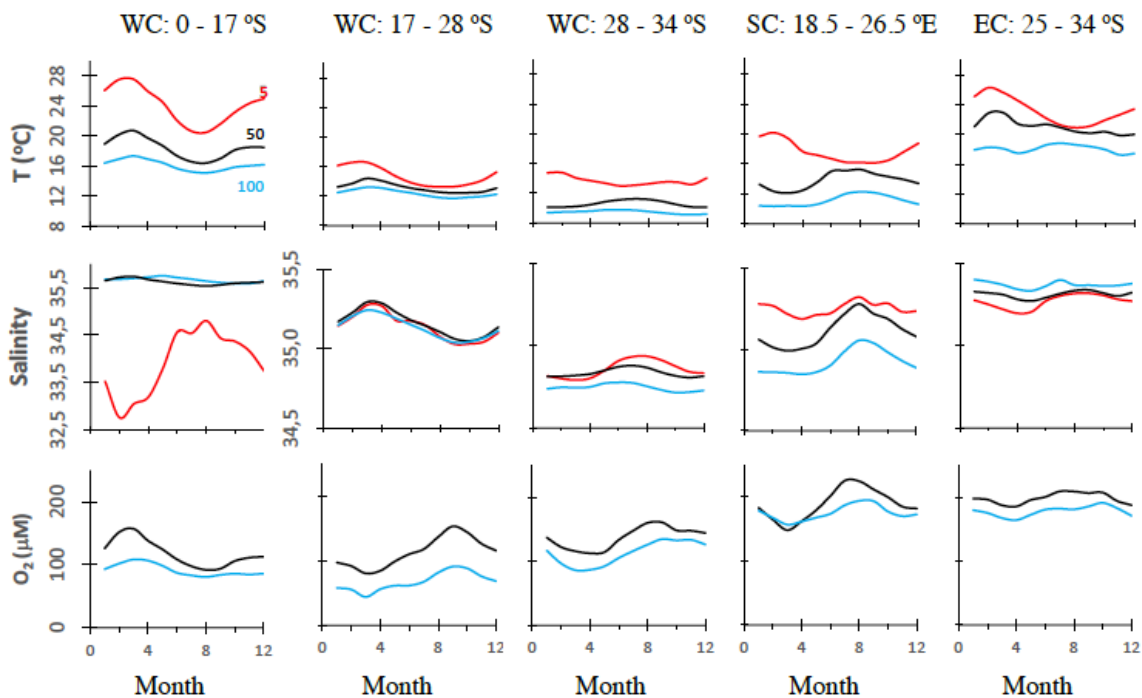


Fig 5. Composite temperature, salinity and oxygen annual profiles at depths of 5 m (red), 50 m (grey) and 100 m (blue), for all the  $H_{\text{coast}}$  bins within 5 degree sections.

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4**Table 1** Percentage of grid boxes requiring gap-filling ( $n < 3$ ), for calculation of temperature at 50 m depth.

AREA	H <sub>coast</sub> Monthly %	Season	H <sub>coast</sub> - H <sub>1.5°</sub> %	H <sub>2°</sub> - H <sub>3°</sub> %	Q1-Q6 %
West Coast: 0 - 17°S	7	JFM	10	51	29
		AMJ	6	32	29
		JAS	2	36	22
		OND	0	17	19
West Coast: 17 - 28°S	0	JFM	0	11	1
		AMJ	0	38	6
		JAS	2	41	6
		OND	2	27	7
West Coast: 28 - 34.5°S	4	JFM	0	14	0
		AMJ	2	33	2
		JAS	0	21	6
		OND	2	33	8
South Coast: 18.5 - 26.5°E	5	JFM	8	23	16
		AMJ	4	21	7
		JAS	2	26	10
		OND	4	17	6
East Coast: 27 - 34.5°S	8	JFM	19	> 50	56
		AMJ	15	> 50	30
		JAS	13	> 50	35
		OND	25	> 50	51
East Coast: 10 - 27°S	25	JFM	13	> 50	> 50
		AMJ	17	> 50	> 50
		JAS	17	> 50	> 50
		OND	36	> 50	> 50
East Coast: 0 - 10°S	33	JFM	18	> 50	> 50
		AMJ	20	> 50	> 50
		JAS	12	> 50	> 50
		OND	23	> 50	> 50

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### 2.3. Analysis of potential sampling bias related to time of day or inter-annual trends

8 Evaluation of data distribution per time of day showed that 20% of the total casts were taken  
9 between 00:00 and 05:59 in the early morning, 32% between 06:00 and 11:59, 28% between 12:00 and  
10 17:59 and 20% between 18:00 and 23:59 (Appendix 1). This fairly even distribution of sampling during  
11 the 24 hour cycle was further explored for potential bias, as follows: a diurnal surface temperature  
12 profile was constructed for each season, consisting of twelve 2-hour average values, for every H bin  
13 with a total of more than 400 casts (Appendix 3). The temperature average obtained from the diurnal  
14 profile ( $T_{\text{diurnal}}$ ) was then compared to the average of all the values in the bin ( $T_{\text{all}}$ ), for each season,  
15 without consideration of sampling time, and yielded the following relationship:

$$16 \quad T_{\text{all}} = 0.9948 * T_{\text{diurnal}} + 0.0827 \quad (r^2 = 0.9988), \text{ and}$$

$$17 \quad (T_{\text{all}} - T_{\text{diurnal}})_{\text{average}} = - 0.003^{\circ}\text{C} \pm 0.105 (1\sigma).$$

18 This analysis showed that there is no appreciable sampling bias related to the time of day, for surface  
19 temperature. It is also worth noting that the 12:00 to 17:59 period was the warmest time of the day (at  
20 the surface) in only 42.5% of the cases evaluated (Appendix 3). It is assumed that the time of day can  
21 also be ruled out as a potential appreciable source of sampling bias at 50 and 100 m depths, and for  
22 salinity and oxygen.

1 The uneven spatial and temporal distribution of data across the study area (Fig 1A; Appendix 1)  
 2 presents a challenge to the analysis of the data for potential inter-annual or inter-decadal trends or  
 3 periodicity. Even in the relatively small geographic area with the largest number of casts over the  
 4 longest period of time, i.e. between 32 and 34°S on the West Coast, time series consisting of one or  
 5 two year average values revealed no statistically significant monotonic trends or periodicity in temperature,  
 6 salinity or oxygen. As an alternative, time series consisting of decadal average values were  
 7 constructed, for 1945-1953 to 2005-2014, for each season and depth interval. A Mann-Kendall test  
 8 was then performed on each constructed time series that comprised of at least 4 decadal values, as a  
 9 test for monotonic trends. Only 3.5% of the H<sub>coast</sub> to H<sub>3°</sub> bins contained enough data, over a long enough  
 10 period and without too many sampling gaps, to satisfy this requirement. All significant monotonic trends  
 11 (>95% confidence level) are listed in Table 2 and the relevance of these trends are discussed in the  
 12 text. All of the East Coast H bins north of 33°S contained inadequate data for inclusion in this analysis.  
 13

14 Although the available data was too sparse for the robust statistical analysis of periodicity using  
 15 spectral analysis, the decadal time series data suggested potentially interesting periodicities in the data,  
 16 generally consistent with the Atlantic multidecadal oscillation (AMO), i.e. alternating periods of warm  
 17 and cold sea-surface temperatures with a period of 60 to 90 years (Kerr, 2000; Knudsen et al., 2011).  
 18 This motivated the construction of three 25 year climatologies (detailed in the next sub-section), for the  
 19 periods 1945 to 1969, 1970 to 1994 and 1995 to 2019, as an alternative evaluation of long-term  
 20 variability. These 25 year periods approximate the timing of one cold (1970-1994) and two warm (1945-  
 21 1969 and 1995-2019) phases of the AMO cycle. It also conveniently divided the data for the West and  
 22 South coast areas into three equal 25 year periods, each with adequate data for the construction of  
 23 climatologies at depths of 5, 50 and 100 m.  
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27 **Table 2** Results of Mann-Kendall tests on decadal time series data for temperature and salinity. Trends  
 28 shown are significant at the >95% level, positive values (bold) indicate increases over time and negative  
 29 values decreasing trends.  
 30

Long (± 0,25°)	Lat	Type	Season	Data Period	T- 5	T-50	T-100	S-5	S-50	S-100
					°C/year			units/year		
11,25	-5,25	H2	JFM	1960-1988	-0,014		-0,026			
13,25	-12,25	H2	JFM	1961-1999	<b>0,050</b>					
12,25	-20,25	H3	JFM	1950-2002	<b>0,022</b>					
14,75	-25,75	H1	JFM	1966-1999	<b>0,033</b>	<b>0,037</b>		<b>0,007</b>	<b>0,005</b>	<b>0,007</b>
18,25	-33,75	H1	JFM	1951-2008			<b>0,016</b>			<b>0,002</b>
19,25	-34,75	H2	JFM	1948-2007		<b>0,079</b>			<b>0,006</b>	
19,75	-34,75	H1	JFM	1948-2007		<b>0,127</b>		<b>0,006</b>	<b>0,010</b>	
12,75	-20,75	H2	AMJ	1961-1999	<b>0,026</b>					
13,75	-23,25	H2	AMJ	1961-2018	<b>0,033</b>	<b>0,020</b>	<b>0,017</b>		<b>0,002</b>	
13,75	-23,75	H2	AMJ	1968-2018	<b>0,076</b>	<b>0,039</b>			<b>0,003</b>	<b>0,003</b>
17,25	-32,25	H3	AMJ	1948-2010	-0,027	-0,022	-0,016		-0,002	-0,002
17,75	-34,25	H2	AMJ	1954-2008	<b>0,049</b>	<b>0,035</b>	<b>0,030</b>		<b>0,003</b>	
23,25	-34,25	H1	AMJ	1958-2008	-0,040					
25,25	-34,25	H1	AMJ	1958-2008	-0,023					
13,25	-21,25	H2	JAS	1959-1998	<b>0,029</b>		<b>0,020</b>		<b>0,002</b>	<b>0,002</b>
16,75	-32,75	H4	JAS	1950-2011	<b>0,016</b>	<b>0,013</b>				
18,25	-33,75	H1	JAS	1950-2008	<b>0,027</b>	<b>0,061</b>	<b>0,020</b>			<b>0,002</b>
17,75	-33,75	H2	JAS	1949-2004	<b>0,022</b>	<b>0,053</b>	<b>0,034</b>	<b>0,003</b>	<b>0,005</b>	<b>0,004</b>
17,75	-34,25	H2	JAS	1950-2008	<b>0,020</b>				<b>0,002</b>	
13,25	-23,25	H3	OND	1964-1999	-0,039			<b>0,003</b>		
14,25	-23,75	H1	OND	1945-1999	-0,018					
19,25	-35,25	H3	OND	1948-2008			<b>0,024</b>			

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#### 2.4. Construction of seasonal and annual climatologies

The calculation of seasonal average values for H bins reduced the amount of gap-filling required considerably, compared to that required for monthly values (Table 1). The lower temporal sampling resolution (that is seasonal as opposed to monthly) then allowed for increased spatial resolution within the following data rich H bins, through their subdivision into  $0.25^\circ \times 0.25^\circ$  (or Q) bins:  $H_{\text{coast}}$  and  $H_{1^\circ}$  along the Gabon-Angola shelf,  $H_{\text{coast}}$  to  $H_{1.5^\circ}$  along the Central and Southern Benguela shelf and the Agulhas Bank, and  $H_{\text{coast}}$  along the East Coast, from  $34.5$  to  $28^\circ\text{S}$ . Along the West and South coasts, seasonal average values were calculated for each of the three variables (T, S and  $\text{O}_2$ ) at depths of 5, 50 and 100 m, for the Q bins inside these H bins, and for the appropriate  $H_{\text{coast}}$  to  $H_{3^\circ}$  bins along the shelf areas outside of these H bins. Along the East Coast, seasonal values were calculated for temperature only. The mixed spatial resolution grid used to construct climatologies took cognizance of the variable continental shelf width of the study area, data density and the observation that in most areas data availability decreases substantially offshore of the 500 m isobath. For gap filling purposes, mean values for "empty" data bins ( $n < 3$ ) were calculated from the four nearest data bins, the pair of bins to the west and east, and that to the north and south. Along the West and South coasts "empty" data bins were usually isolated from each other and gap filling was straightforward. Along the East Coast, however, the larger proportion of "empty" data bins (Table 1) usually required an iterative process, starting with the filling of "empty" bins surrounded by at least 3 or 4 bins containing data, followed by the filling of empty bins surrounded by 2 bins containing data in addition to perhaps 1 or 2 gap-filled bins, and lastly empty bins surrounded by at least 1 data bin and 2 to 3 gap-filled data bins.

After construction of the seasonal climatologies, based on all data from 1945 to 2019 (Fig 6, 7, 8), annual climatologies consisting of the average of the seasonal values were constructed (Fig 9). Along the East Coast, the annual climatologies for salinity and oxygen represent the average of all available data. This procedure was then repeated to construct annual climatologies for the West and South coast areas, representing the 1945-1969, 1970-1994 and 1995-2019 periods mentioned earlier, but at a  $0.5^\circ$  scale, i.e. using H bins only. Differentials were calculated from the 25 year climatological averages, and mapped as follows for temperature, (Fig 10), and similarly for salinity (Fig 11) and oxygen (Fig 12).

A to C:  $\text{Avg}_{[1970 \text{ to } 1994]} - \text{Avg}_{[1945 \text{ to } 1969]}$

D to F:  $\text{Avg}_{[1995 \text{ to } 2019]} - \text{Avg}_{[1970 \text{ to } 1994]}$

G to I:  $\text{Avg}_{[1995 \text{ to } 2019]} - \text{Avg}_{[1945 \text{ to } 1969]}$

J to L:  $\text{Avg}_{[1995 \text{ to } 2019]} - \text{Avg}_{[1945 \text{ to } 1969]}\#$ , with # indicating the exclusive plotting of H bins for which 25 year climatological values are available for each of the three 25 year periods, and that display monotonic trends (increasing or decreasing) from 1945 to 2019.

The standard deviation associated with the average of all of the data (1945 to 2019) within each H or Q was mapped, as well as the standard deviation associated with the average of the seasonal values (Appendix 4). These statistics represent different types of variability within the data set, i.e. the total variability in the data and the variability around the seasonal average, and are not equivalent. All climatologies were plotted using the ODV free software package (Schlitzer, 2018). The "quick gridding"

1 display option was used, since it produced the most realistic and least distorted data displays. The  
2 construction and visualization of the climatologies presented in this study are purposefully simplistic, in  
3 order to encourage its use, reproduction and modification.  
4

### 5 **3. INTRA-ANNUAL VARIABILITY ALONG THE SOUTHERN AFRICA CONTINENTAL SHELF**

#### 6 *3.1. Gabon to Angolan shelf, from the equator to ~ 17°S on the West Coast*

7 Along the southern Africa shelf, the largest seasonal surface temperature fluctuations are observed on  
8 the shelf area between Gabon and Angola (or the Northern Benguela) (Fig 3A, 5). It is also only along  
9 the Gabon-Angola and the Central Benguela shelf areas that similar seasonal temperature profiles are  
10 observed at depths of 5, 50 and 100 m (Fig 5). These two adjacent areas also display similar  
11 temperature seasonality, or occurrence of the warmest and coldest periods of the annual cycle, and at  
12 all depths (Fig 4). During summer, the shelf area north of ~ 12°S is characterized by water temperatures  
13 exceeding 28°C, and very strong temperature gradients are present south of that towards the 20°C  
14 isotherm (Fig 3A, 6A). Apart from this, > 28°C surface water temperatures are observed on the shelf  
15 only in autumn (Fig 6B). The inner Angolan shelf area is dominated by 24 to 26°C surface water in  
16 autumn and spring, and 20 to 22°C in winter (Fig 6C, 6D). At 50 m depth an intrusion of warm water  
17 (> 20°C) is evident along the Gabon-Angola shelf, that appears to initiate near the equator in spring,  
18 and penetrates as far south as 12°S in summer, before it retreats or dissipates in autumn (Fig 6E-H).  
19 A similar seasonal intrusion of warm water (> 16°C) is observed at 100 m (Fig 6I-L). The seasonal  
20 incursion of warmer water observed in the climatologies, supports the proposed strengthening of the  
21 Gabon-Congo Undercurrent (GCUC) during spring (Stramma and Schott, 1999). The GCUC flows  
22 southwards along the African coast and derives primarily from the Equatorial Under Current between  
23 1°S and 6°S (Stramma and Schott, 1999).

24 Surface water on the Gabon-Angola shelf is significantly influenced by fresh water outflow from  
25 several large rivers, most notably the Congo River, the largest river along the southern Africa coastline  
26 (Fig 1A, 5). This results in large seasonal variability, and pronounced salinity gradients at the surface  
27 and between the surface and water at 50 and 100 m (Fig 5). The surface salinity climatologies show  
28 that the Congo river plume influences a very large part of the shelf in summer (Fig 7A) and reaches its  
29 smallest extent in winter (Fig 7C). This is consistent with Congo river seasonal flow fluctuations (Sonwa  
30 et al., 2020). Similar seasonal changes in salinity, but of smaller magnitude, just south of the equator  
31 and at ~8°S, presumably reflect the influences of the Ogooué river in Gabon and the Cuanza river in  
32 Angola, respectively.

33 The Gabon-Angola shelf area is also characterized by pronounced spatial and seasonal changes  
34 in dissolved oxygen, particularly at depths of 50 and 100 m (Fig 5, 8E-L). The oxygen climatologies  
35 show generally decreasing values from the equator towards the Angolan shelf. At 100 m depths, low  
36 oxygen water (< 100 µM O<sub>2</sub>) water dominates the shelf between 5°S and 17°S throughout the year (Fig  
37 8I-L). At both 50 and 100 m depth there is a westward and northward expansion of < 100 µM O<sub>2</sub> water  
38 from summer (Fig 8E, 8I) to winter (Fig 8G, 8K). During summer, < 100 µM O<sub>2</sub> water is situated offshore  
39 of the Angola shelf at 50 m depth (Fig 8E), as is the case for < 50 µM O<sub>2</sub> water at 100 m depth (Fig 8I).  
40 This climatological observation of a minimum in the westward and northward extension of low oxygen  
41 water in summer, and a maximum extension of this water in winter, is consistent with previous reports  
42 for this area (Chapman and Shannon, 1987; Mohrholz et al., 2008).

1

### 2 3.2 Central and Southern Benguela shelf, from ~ 17°S to 35°S on the West Coast

3 Along the southern African shelf area, the smallest seasonal temperature ranges, at all depths, are  
4 observed in the climatologies of the Southern Benguela (28 to 36°S) inner shelf, and the Central  
5 Benguela (17 to 28°S; Fig 3, 5). At ~ 27°S and 32°S, where the shelf is at its broadest (Fig 1A), summer  
6 surface temperatures are higher and the seasonal temperature range larger, compared to adjacent  
7 shelf areas (Fig 3A). A notable feature of seasonal temperature variability in the Southern Benguela is  
8 that temperature moves out of phase with increasing depth (Fig 4, 5). At 50 and 100 m depths,  
9 temperature is higher during the winter than during spring and summer, the opposite to what is observed  
10 at the surface. This result is consistent with the seasonal Southern Benguela upwelling season, which  
11 peaks in spring and summer (Hutchings et al., 2009) and is associated with the upwelling of colder (and  
12 lower salinity) water onto the shelf. Upwelling processes in the Central and Southern Benguela have  
13 been extensively studied and there are numerous publications and reports that can be consulted for  
14 much more detail than the space in this manuscript allows for ( Hutchings et al., 2009, and reference  
15 therein; Jarre et al., 2015b; Lamont et al., 2018).

16 Seasonal climatologies show that surface temperature values below 16°C (as an example) is a  
17 persistent feature along the inner shelf along the Central and Southern Benguela in spring and autumn  
18 (Fig 6B, 6D). During summer, however, this 16°C surface isotherm contracts towards the coast and is  
19 absent from the northern half of the Central Benguela (Fig 6A). During winter, in contrast, < 16°C water  
20 extends much further westward, as well as further north and south than is observed in spring and  
21 autumn (Fig 6C). During winter, in fact, the surface 16°C isotherm extends beyond the Southern  
22 Benguela, in an eastward direction along the shelf to ~ 22°E on the Agulhas Bank (Fig 6C). Similarly,  
23 16 to 18°C surface water is present in a continuous band that extends from the Central Benguela shelf  
24 to the eastern Agulhas Bank in spring and autumn (Fig 6B, 6D). In contrast, 16 to 18°C surface water  
25 is restricted to the Central and Southern Benguela and absent from the Agulhas Bank in summer, (Fig  
26 6A), but absent from the Central and Southern Benguela, and present on the Agulhas Bank, in winter  
27 (Fig 6C). These seasonal contractions and expansions in the areal extent of water within specific  
28 temperature envelopes are also observed in the 50 and 100 m climatologies. At 100 m depth, for  
29 example, < 12°C is present on the shelf in a continuous band that extends from the Southern Benguela,  
30 all the way to the eastern edge of the Agulhas Bank, in summer and autumn (Fig 6I, 6J). In winter (Fig  
31 6L) and spring (Fig 6L), however, the < 12°C extends into the Central Benguela, at a depth of 100 m.  
32 Also, in winter and spring, < 12°C water does not occur in a continuous band from the Southern Benguela  
33 to the eastern Agulhas Bank at 100 m, due to the presence of warmer water between ~ 20 and 22°E  
34 during these seasons.

35 In the Central Benguela, seasonal salinity climatologies (for 5, 50 and 100 m) suggest a southward  
36 shift in the position of the salinity isolines in the northern part of the Central Benguela shelf (~ 17 to  
37 22°S, Fig 7), consistent with the southward migration of the ABF, and the higher salinity and low oxygen  
38 water associated with it, from February to April (Dias, 1983; Stramma and Schott, 1999; Ekau and  
39 Verheye, 2005; Mohrholz et al., 2008). The Southern Benguela shelf is characterized by the lowest  
40 salinity waters on the southern African shelf, with the obvious exception of surface water on the Gabon-  
41 Angolan shelf (Fig 5). The surface salinity climatologies show a freshwater influence at the Orange  
42 river mouth (~28.5°S) and at the Berg river mouth (~ 32.5°S) (Fig 7A-D). This is more pronounced in

1 summer and winter respectively, consistent with the peak runoff seasons of these rivers (de Villiers and  
2 Thiart, 2007). Apart from these river influences, the lowest salinity waters are found at 100 m depth,  
3 and during summer (Fig 7I) and autumn (Fig 7J). Enhanced freshening of water during summer and  
4 autumn are also observed at 50 m at  $\sim 32^{\circ}\text{S}$  in the Southern Benguela. The seasonal salinity changes  
5 evident in the 50 and 100 m climatologies, are consistent with the seasonal upwelling of low salinity  
6 water onto the shelf (Hutchings et al., 2009).

7 Composite graphs of intra-annual variability in dissolved oxygen show that the Central  
8 Benguela is characterized by the lowest oxygen levels in the study area (Fig 5). In both the Central  
9 and Southern Benguela, oxygen levels decline during the summer and autumn, at both 50 and 100 m  
10 depths, and reach a seasonal high in winter. The seasonal oxygen climatologies show that the  
11 presence of low oxygen water ( $< 100 \mu\text{M O}_2$ ) is widespread along the inner shelf at 100 m depth (Fig  
12 8I-L) in both the Central and Southern Benguela. The areal extent of this low oxygen water is more  
13 widespread in summer and autumn, than in winter and spring. In the Central Benguela, these seasonal  
14 changes are more pronounced, and low oxygen water extends  $\sim 2^{\circ}$  further south and further offshore  
15 during summer and autumn. These seasonal changes are consistent with the seasonal incursion of low  
16 oxygen water from the Angola Gyre into the Central Benguela, mentioned earlier (Dias, 1983;  
17 Stramma and Schott, 1999; Ekau and Verheye, 2005; Mohrholz et al., 2008) and documented  
18 seasonality in the Southern Benguela (Hutchings et al., 2009; Jarre et al., 2015b). At 50 m depth, low  
19 oxygen water is less widespread than at 100 m (Fig 8E-H). In the Southern Benguela, low oxygen water  
20 is evident at 50 m in the summer, and at the  $\sim 29^{\circ}\text{S}$  and  $\sim 32.5^{\circ}\text{S}$  upwelling cells only (Fig 8E). There  
21 is some indication of its presence, at 50 m, at the  $\sim 29^{\circ}\text{S}$  upwelling cell in autumn (Fig 8F) and the  $\sim$   
22  $32.5^{\circ}\text{S}$  cell in spring (Fig 8H), but it is entirely absent from the Southern Benguela in winter (Fig 8G).  
23 Surface ocean oxygen levels (Fig 8A-D) vary seasonally in response to primarily changes in primary  
24 productivity and temperature, and a discussion of these complex processes and interpretation of these  
25 changes are beyond the scope of this manuscript.

### 26 27 *3.3 South Coast or Agulhas Bank, from $18.5^{\circ}\text{E}$ to $26.5^{\circ}\text{E}$*

28 A prominent commonality between the South Coast and the Southern Benguela, is the contrasting  
29 temperature seasonalities of surface and deeper waters (Fig 5). On the Agulhas Bank, as also  
30 observed in the Southern Benguela, water temperature at 50 and 100 m depths reaches a seasonal  
31 high in winter, and not in summer as is observed at the surface. On the Agulhas Bank, this difference  
32 in seasonality between surface and deeper waters is more pronounced than in the Southern Benguela.  
33 Temperature climatologies show that along the South Coast, surface temperature generally increases  
34 in an easterly direction (Fig 6A-D). Along the inner shelf, this general trend is disrupted in areas of  
35 coastal upwelling, for example between  $\sim 23^{\circ}$  and  $24^{\circ}\text{E}$  in autumn (Fig 6B). During summer, the  
36 Agulhas Bank is dominated by surface water temperatures in the  $20$  to  $22^{\circ}\text{C}$  range (Fig 6A). From  
37 autumn to spring, however, the  $20^{\circ}\text{C}$  surface isotherm is situated along the eastern shelf edge of the  
38 Agulhas Bank, and the shelf is dominated by  $16$  to  $20^{\circ}\text{C}$  surface water (Fig 6B-D). The  $16^{\circ}\text{C}$  surface  
39 isotherm migrates from its spring/summer inner shelf position at Cape Point ( $\sim 18.5^{\circ}\text{E}$ ) to just west of  
40 Cape Agulhas ( $\sim 20^{\circ}\text{E}$ ) in autumn and Cape Infanta ( $\sim 21^{\circ}\text{E}$ ) in winter (Fig 6C). At 100 m depth, cold  
41 ( $< 16^{\circ}\text{C}$ ) water is present along the inner shelf throughout the year, but it is more prominent and present  
42 in a continuous band than extends into the Southern Benguela, in summer (Fig 6I) and autumn (Fig 6J),

1 as mentioned previously. Temperature climatologies suggest similar seasonalities are present at 50 m  
2 (Fig 6E-H).

3 Seasonal salinity climatologies (Fig 7), mirror that of temperature, with lower salinities coinciding  
4 with colder temperatures, and higher salinities with higher temperatures. Seasonal oxygen  
5 climatologies (Fig 8) show that cold, lower salinity waters also typically have lower oxygen values, with  
6 similar seasonal variations to that observed for temperature and oxygen. Although Agulhas Bank water  
7 has higher O<sub>2</sub> levels than that of the Southern Benguela shelf, ranging from 125 to 200 μM at 100 m  
8 compared to 75 to 175 μM, it displays similar seasonality, with generally lower values in summer and  
9 autumn and to a lesser extent spring. The upwelling features and gradients across the Agulhas Bank,  
10 evident in the temperature, salinity and oxygen climatologies are consistent with previous observations  
11 of wind-driven coastal upwelling and shelf-edge upwelling induced by the southward flowing Agulhas  
12 Current (Probyn et al., 1994; Lutjeharms et al., 1996). The seasonal climatologies presented here  
13 provide new insight into the spatial extent of intra-annual variability and the connectivity with the  
14 Southern Benguela.

15

### 16 *3.4 East Coast, from the equator to 36° S*

17 Despite the limitations resulting from relative data scarcity along the East Coast, some patterns are  
18 evident. Along the East Coast, there is a decrease in temperature, but a general increase in the  
19 seasonal surface temperature range from the equator towards the pole (Fig 3), which is opposite to the  
20 trend observed along the West Coast. Seasonal temperature profiles appear complex, particularly  
21 along the inner shelf and at 100 m depth (Fig 5). This may be an artefact of the relative scarcity of  
22 data on the East Coast, but it can also reflect high variability associated with the incursion of offshore  
23 waters onto the shelf, which is very narrow in some areas compared to the West and South Coast shelf  
24 areas. East Coast shelf water is consistently warmer than that of the West Coast at similar latitudes  
25 (Fig 3). For example, between 0 and 10°S, East Coast shelf water is approximately 6 to 8°C warmer  
26 than that of the West Coast, at all three depths. Between 25 and 35°S, this offset is 8 to 10°C at each  
27 of the three different depths.

28 It is evident in the seasonal surface climatologies that > 28°C water is much more common on the  
29 the East Coast, than on the West Coast, and it is present as far south as at least 25°S in summer (Fig  
30 6A) and 15°S in autumn and spring (Fig 6B, 6D). In winter, however, the East Coast shelf is  
31 characterized by average surface temperatures below 26°C (Fig 6C). The 24°C surface isotherm  
32 intersects the coastline at a relatively narrow latitudinal range of 25 to 27°S from autumn to spring, but  
33 extends much further south in summer. This southward summer shift coincides with the westward shift  
34 of the 20°C isotherm along the southern margin of the East Coast and the Agulhas Bank, and an  
35 apparent shelf-ward shift of offshore warm surface water during this season, possibly related to  
36 seasonality in the position of the Agulhas Current. An interesting feature of the temperature  
37 climatologies is the temperature maximum between 10 and 15°S along the East Coast, particularly  
38 noticeable at deeper depths (Fig 6D-H). The warmer shelf water between these latitudes is also  
39 characterized by lower salinity values, compared to shelf water further north and south.

40 Freshening of surface water is evident offshore of river mouths, for example the Tugela, Limpopo,  
41 Zambezi, Rufiji and Tana rivers (Fig 1, Fig7A-D). The geographic extent of the influence of these river  
42 outflows on the salinity of shelf surface water, however, is not as pronounced as that of the Congo river

1 on the Angolan shelf area. Data scarcity impedes capture of such features in a climatology. The warm  
2 waters of the East Coast contain high oxygen levels with apparently low seasonal variability; the scarcity  
3 of oxygen data along the East Coast, however, precludes a more detailed discussion and evaluation of  
4 these observations.

#### 6 **4. ENVIRONMENTAL CLIMATOLOGIES AND ECOLOGICAL CONSIDERATIONS**

##### 7 *4.1 Biogeographic zone boundaries in comparison to key climatological map features*

8 The approximate boundaries of biogeographic Realms and ecoregions blocks (Spalding et al., 2007),  
9 LME's (Sherman, 1993) and coastal biogeographic zones (Potts et al., 2015), were superimposed on  
10 the annual climatologies of temperature, salinity and oxygen, to qualitative evaluate biogeographic  
11 boundaries in the context of environmental spatial patterns (Fig 9). It is difficult to find qualitative  
12 agreement between the temperature, salinity and oxygen climatologies, and the large scale  
13 biogeographic zones, that is the Realms and LME's. If the Temperature Southern African Realm (Fig  
14 2) extended  $\sim 5^\circ$  further north, with its boundary with the Tropical Atlantic Realm positioned just north  
15 of the ABF, it would capture all southern African shelf water with average surface temperatures  $< 24^\circ\text{C}$ .  
16 The smaller ecoregion blocks within these Realms (numbered in Fig 9 according to Spalding et al.,  
17 2007), however, also display poor qualitative agreement with environmental climatologies. An  
18 exception to this is the "Namaqua" ecoregion block (#191 in Fig 2 and 9), which corresponds to the  
19 Southern Benguela (Fig 9). The LME divisions (Sherman, 1993) show the poorest relationship with  
20 environmental variables. The coastal biogeographic zones (Potts et al., 2015), overall, agree the best  
21 with features of the environmental climatologies. This classification scheme uniquely transcends  
22 national boundaries, and also rely primarily on taxonomy (Briggs and Bowen, 2012; Potts et al. 2015;  
23 Whitfield, 2005). The tropical coastal zone corresponds with average surface water temperatures  
24  $> 25^\circ\text{C}$  on both the West and East Coast, the sub-tropical coastal zone with average surface values  
25 between  $\sim 23$  and  $25^\circ\text{C}$  on both the West and East Coast, and the warm-temperate coastal zone with  
26 surface values between  $\sim 17$  and  $23^\circ\text{C}$ , on the West and South Coast (Fig 9A). The construction of  
27 climatologies for key ocean variables, and at the surface as well as deeper depths, can be used as a  
28 guide to refine the boundaries of biogeographic zones, or possibly to evaluate the usefulness of the  
29 concept of such boundaries in a highly variable environment.

##### 31 *4.2 Decadal changes in environmental climatologies, compared to marine species range shifts*

32 Differences between the three 25 year climatologies, representing changes and trends from 1945 to  
33 2019, demonstrate the complexity and spatial heterogeneities inherent in environmental change, that  
34 have been the bugbear of efforts to understand increasing evidence for ecosystem change (Fig 10 to  
35 12). For the sake of brevity, only the most prominent changes are summarized below, followed by a  
36 synopsis of documented environmental and ecosystem changes, that correspond to changes and  
37 trends evident in the environmental climatologies (Table 2; Fig 10-12):

- 38
- 39 • The Gabon-Angola shelf was generally cooler in 1970-1994 compared to both the 1945-1969 and  
40 1995-2019 periods.



- 1 • Temperature change in the Central and Southern Benguela were generally out of phase with that on
- 2 the Gabon-Angola shelf in 1970-1994, compared to 1945-1969, while the Agulhas Bank was in phase
- 3 with the Gabon-Angola shelf during these periods.
- 4 • There is a general warming trend along the entire West Coast shelf area, at all depths; the only
- 5 exceptions to this warming trend are observed at some of the upwelling areas, most notably parts of
- 6 the Angolan shelf and the central part of the Southern Benguela.
- 7 • Warming is more pronounced in the Northern Benguela, than in the Central Benguela.
- 8 • There is a warming trend on the western Agulhas Bank, that contrasts with a cooling trend on the
- 9 eastern Agulhas Bank; these trends are more prominent at deeper depths around the shelf edge.
- 10 • Increasing temperature trends are generally accompanied by increasing trends in salinity, and
- 11 cooling trends with decreasing trends in salinity.
- 12 • Dissolved oxygen levels are generally decreasing in the Central and Southern Benguela, and the
- 13 Agulhas Bank, but there is evidence for increasing oxygen levels in the Northern Benguela.

14

15 The above climatological trends are consistent with documented changes in the environment, such

16 as:

- 17 • Increasing trends in sea surface temperature off Angola over the past three decades, with above
- 18 global average values (Hobday and Pecl, 2014; Bindoff et al., 2019) of between 0.23°C (Jarre et al.,
- 19 2015a) and 0.8°C (Potts et al., 2014) per decade.
- 20 • Southward shifts of warm Angola-Benguela Front water by about 2°, in response to changes in the
- 21 position of the South Atlantic High-Pressure Cell (Vizy et al., 2018).
- 22 • A decline in upwelling favourable winds in the Central Benguela, but increases in the Southern
- 23 Benguela (Hutchings et al., 2009; Jarre et al., 2015; Van der Lingen and Hampton, 2018).
- 24 • A decline of upwelling off Namibia (Santos et al., 2012; Lamont et al., 2018) and coastal temperature
- 25 increases of between 0.2 and 0.5°C per decade in the northern part of the Central Benguela (Jarre et
- 26 al., 2015a).
- 27 • Cooling of inshore waters along South Africa's west and south coast by 0.1 to 0.2°C per decade
- 28 over the past four decades, in contrast to warming of Agulhas Current water by up to 0.6°C per
- 29 decade (Rouault et al., 2010; Blamey et al., 2015).
- 30 • Increased cooling and upwelling in the Southern Benguela (Rouault et al., 2010; Lamont et al.,
- 31 2018; Leduc et al., 2010; Santos et al., 2012).
- 32 • Broadening of the Agulhas Current (Beal and Elipot, 2015). Differences between the 25 year
- 33 climatologies suggest that long-term cooling of the eastern Agulhas Bank is taking place (in areas of
- 34 wind-induced coastal upwelling and along the shelf edge), with a contrasting warming trend on the
- 35 western Agulhas Bank, that is on the "lee" side of the Agulhas Bank in relation to the flow direction of
- 36 the Agulhas Current. The results suggest that the Agulhas Bank may be significantly impacted by the
- 37 reported broadening of the Agulhas Current.

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39 The climatological trends are also consistent with documented ecosystem changes, such as:

- 40 • Decreased pelagic fish catches off Angola from the late 1970's to 1990's, with recoveries since the
- 41 mid-2000's (Kainge et al., 2020). Oxygen has been shown to be a key determinant of species such

1 as juvenile deepwater hake in the Northern Benguela (Kainge et al., 2017) and the climatologies  
2 document warmer temperature and higher oxygen levels, coinciding with the recovery of fisheries.  
3 • An increase in the latitudinal range of demersal species off Angola, with an overall southward  
4 tendency and an expansion into deeper water (Yemane et al., 2014).  
5 • A southward shift in the distribution of some Angolan coastal fish species (Potts et al., 2014)  
6 • A southward shift for round herring in the Southern Benguela (Blamey et al., 2015).  
7 • An eastward shift for adult anchovy (Roy et al., 2007; Fairweather et al., 2006) and sardine (Coetzee  
8 et al., 2008) since the mid-to-late-1990's (Augustyn et al., 2017).  
9 • An eastward migration of rock lobster (Cockcroft et al., 2008).

10  
11 There are several important caveats associated with the choice of 25 year periods for the  
12 construction of climatologies, that need to be emphasized. The choice of periods were primarily  
13 informed by the availability of data, for both the West and South Coast shelf areas. It is not the intention  
14 of this study to propose that these are the best or most ideal periods for the evaluation of decadal  
15 periodicities or long-term trends, or that the entire southern African shelf area is subject to decadal  
16 periodicities with similar amplitudes or periods. The choice of periods may not even be ideal for the  
17 evaluation of Gabon-Angola shelf waters, despite climatological evidence for AMO periodicities on the  
18 adjacent continental shelf and the ocean areas further north (Kerr, 2000; Knudsen et al., 2011; Sonwa  
19 et al., 2020). The most appropriate choice of time period for the study of trends or periodicities can  
20 only be informed by the availability of more data. These caveats, however, make the correspondence  
21 between changes observed in the climatologies, compared to documented and anticipated (Bakun et  
22 al., 2015; Bindoff et al., 2019) environmental and ecological changes, all the more remarkable. A  
23 speculative observation, based on the results of this study, is that decadal periodicities are in the  
24 process of being interrupted by a longer-term trend, possibly as the result of global warming.

## 25 26 **5. CONCLUSIONS**

27 This study raises several important issues in regards to marine species range shifts, in response to  
28 environmental change, and the definition of biogeographic zone boundaries. Ecological changes can  
29 only be fully understood and interpreted in an environmental context, if the environmental tolerances  
30 and ranges of species are known. Environmental climatologies, such as those presented here, can be  
31 used to guide the design of laboratory and field studies, for the optimal collection of data aimed at better  
32 understanding the environmental tolerance limits of species. The climatologies also shed important light  
33 on the geographic complexity of environmental change, that can be used to guide management  
34 decisions about geographic areas to prioritize for monitoring and conservation purposes. The  
35 climatologies also demonstrate that ecoregion or biogeographic boundaries cannot be assumed to  
36 occupy similar geographic locations for species that occupy different depth zones, and that changes  
37 observed at the surface cannot be assumed to apply to deeper depths.

38 Data availability is a concern, most obviously for the East Coast shelf area. There are two specific  
39 areas that appear curiously under-studied and under-sampled, given their oceanographic significance  
40 in the region: the Agulhas Current source region off southern Mozambique on the East Coast, and the  
41 large Lüderitz upwelling at ~ 27°S on the West Coast. The Agulhas Bank has been more extensively  
42 studied than these two areas, but not nearly as extensively as the Southern Benguela; the magnitude

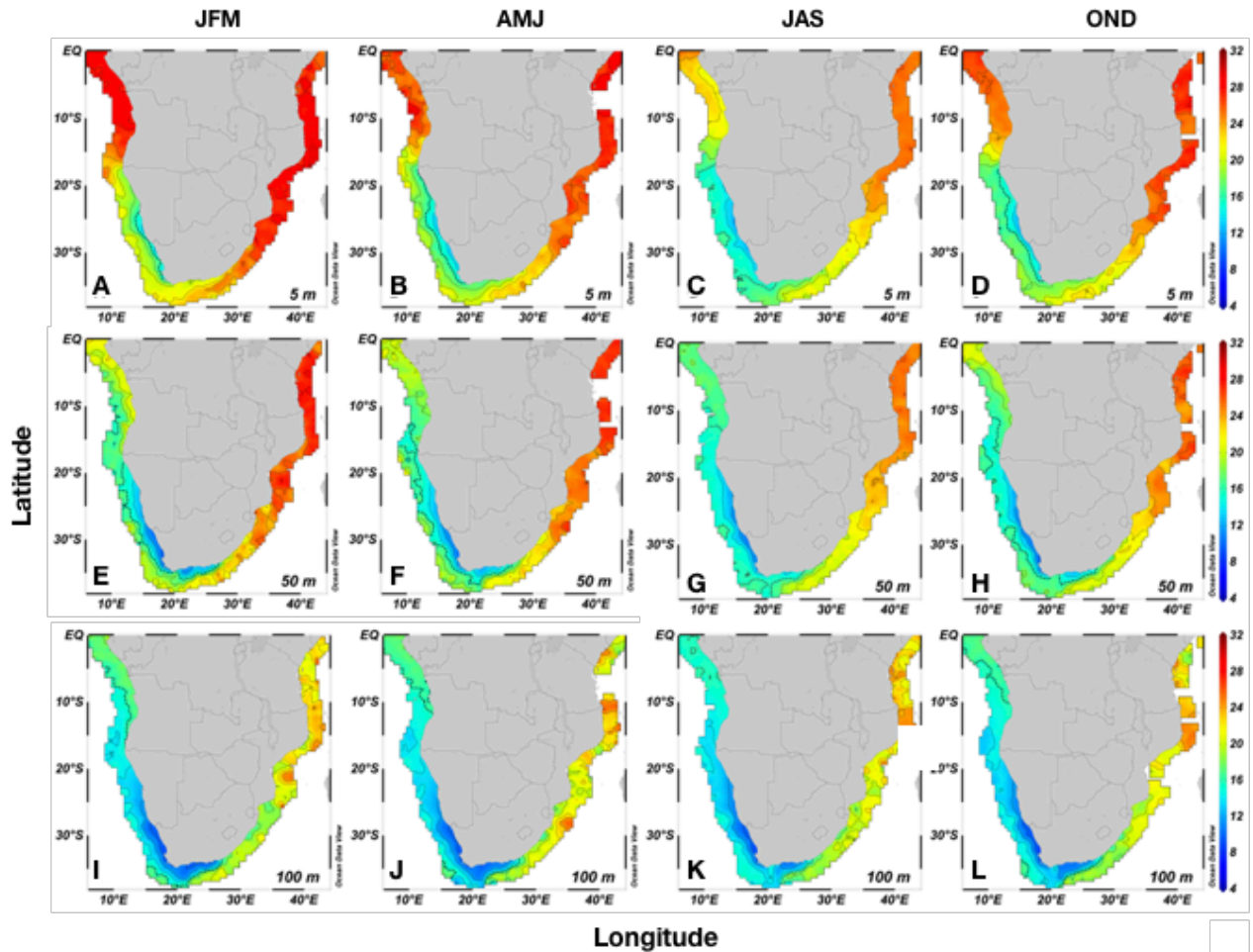
1 of change evident for the Agulhas Bank in the climatologies should motivate more studies and  
2 monitoring in this important transition zone. The most significant concern around data availability,  
3 however, is that only 0.33% of the data used in this study represent sampling carried out since 2010  
4 (Appendix 1). In some instances this reflects an actual reduction in ocean monitoring activities, but it  
5 also reflects changing attitudes towards the placement of data in the public domain, even when data  
6 collection is publicly funded. There is evidence for accelerated global environmental change since  
7 2010, (Bindoff et al., 2019), that is not captured in the climatologies of this study, because of the dearth  
8 of recent data in the public domain for the southern African shelf area. It remains to be seen, therefore,  
9 to what extent the decadal changes presented in this study over- or under-represent present and on-  
10 going change in the region.

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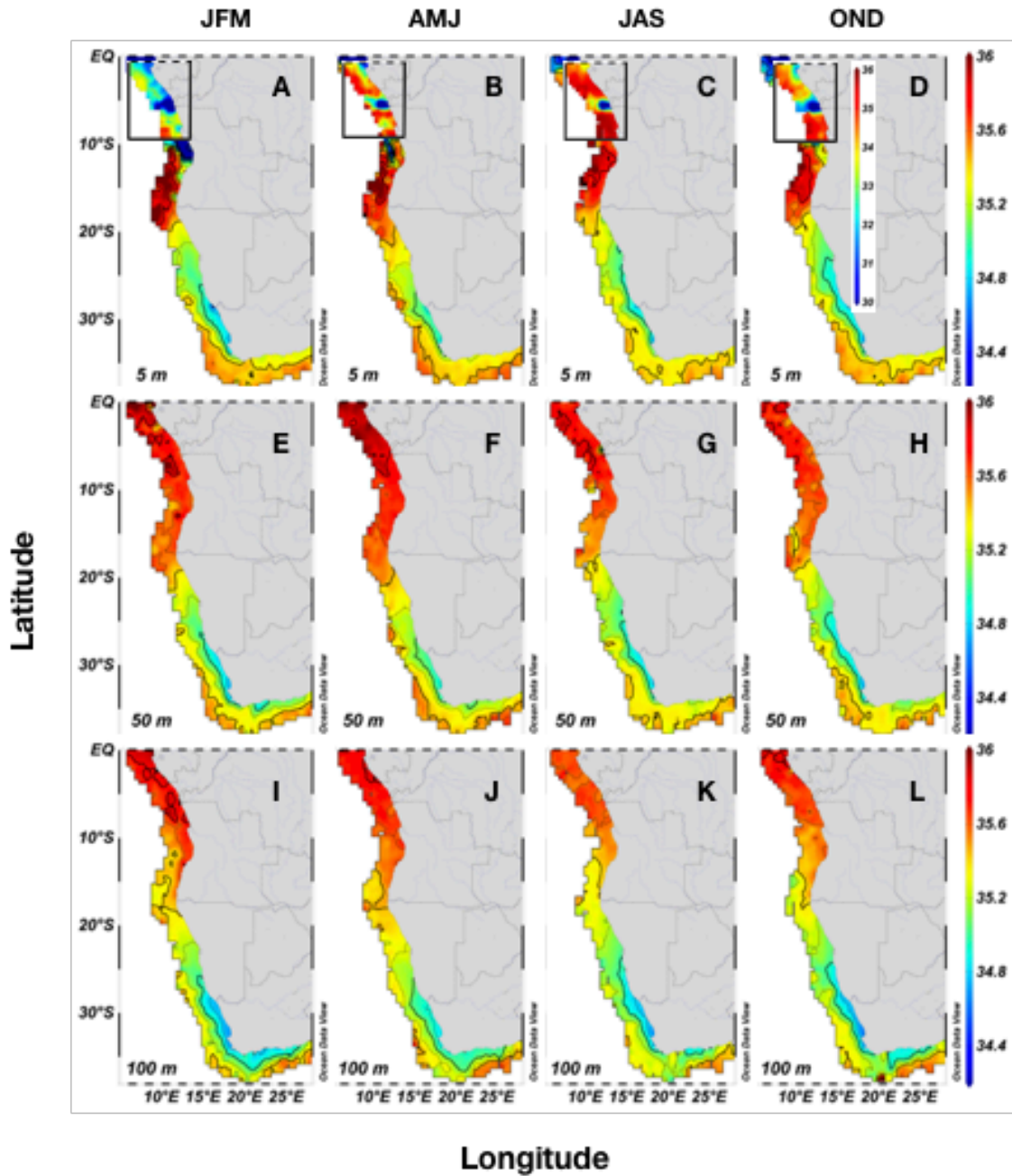
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Fig 6. Seasonal temperature climatologies at depths of 5 (A to D), 50 (E to H) and 100 m (I to L) for summer (JFM), autumn (AMJ), winter (JAS) and spring (OND). Contour intervals are 2°C, with 16°C indicated with a dashed line.

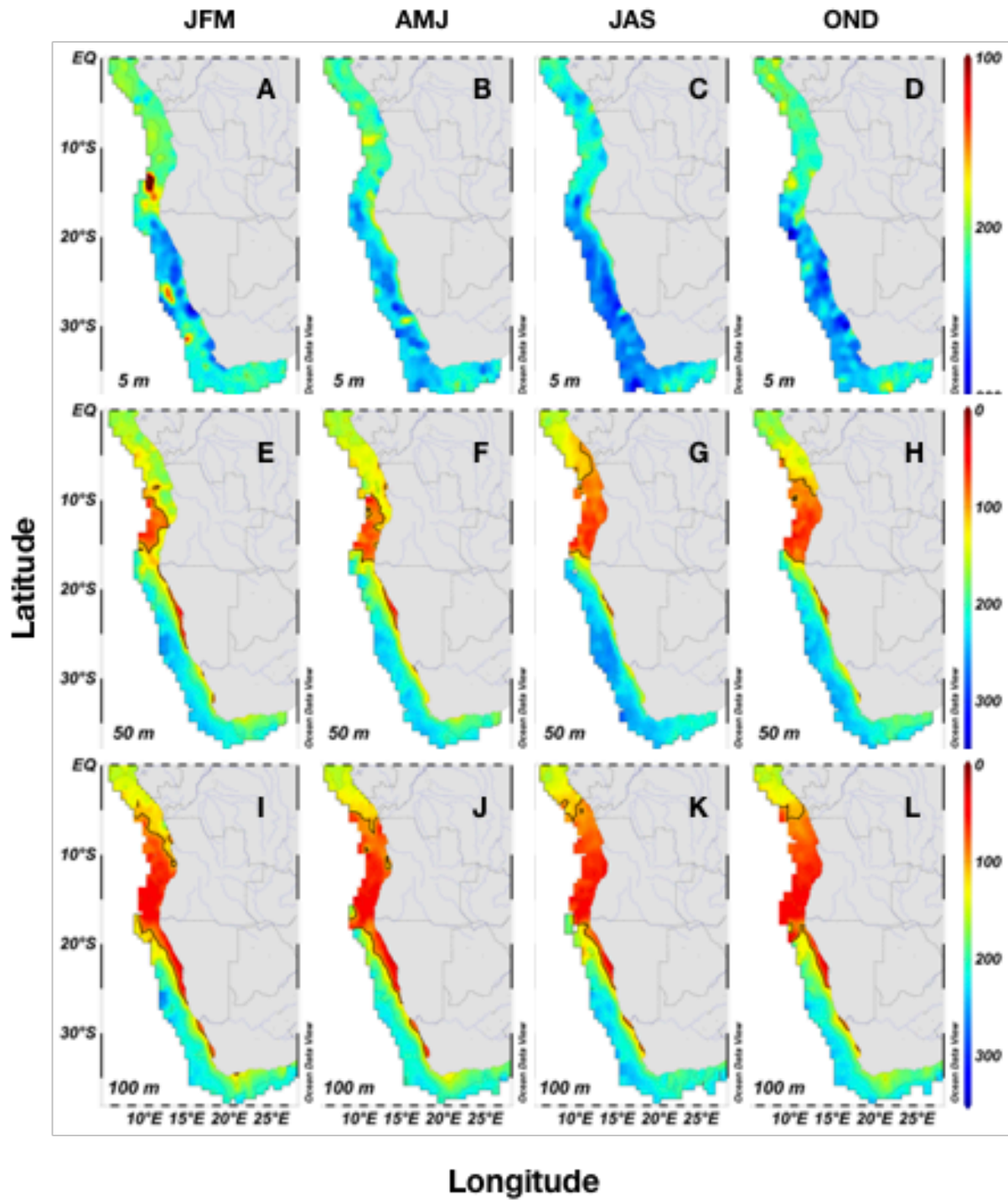
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Fig 7. Seasonal salinity climatologies at depths of 5 (A to D), 50 (E to H) and 100 m (I to L) for summer (JFM), autumn (AMJ), winter (JAS) and spring (OND). Contour intervals are 0.2, with 34.6, 35.0 and 35.4 indicated with thicker lines.

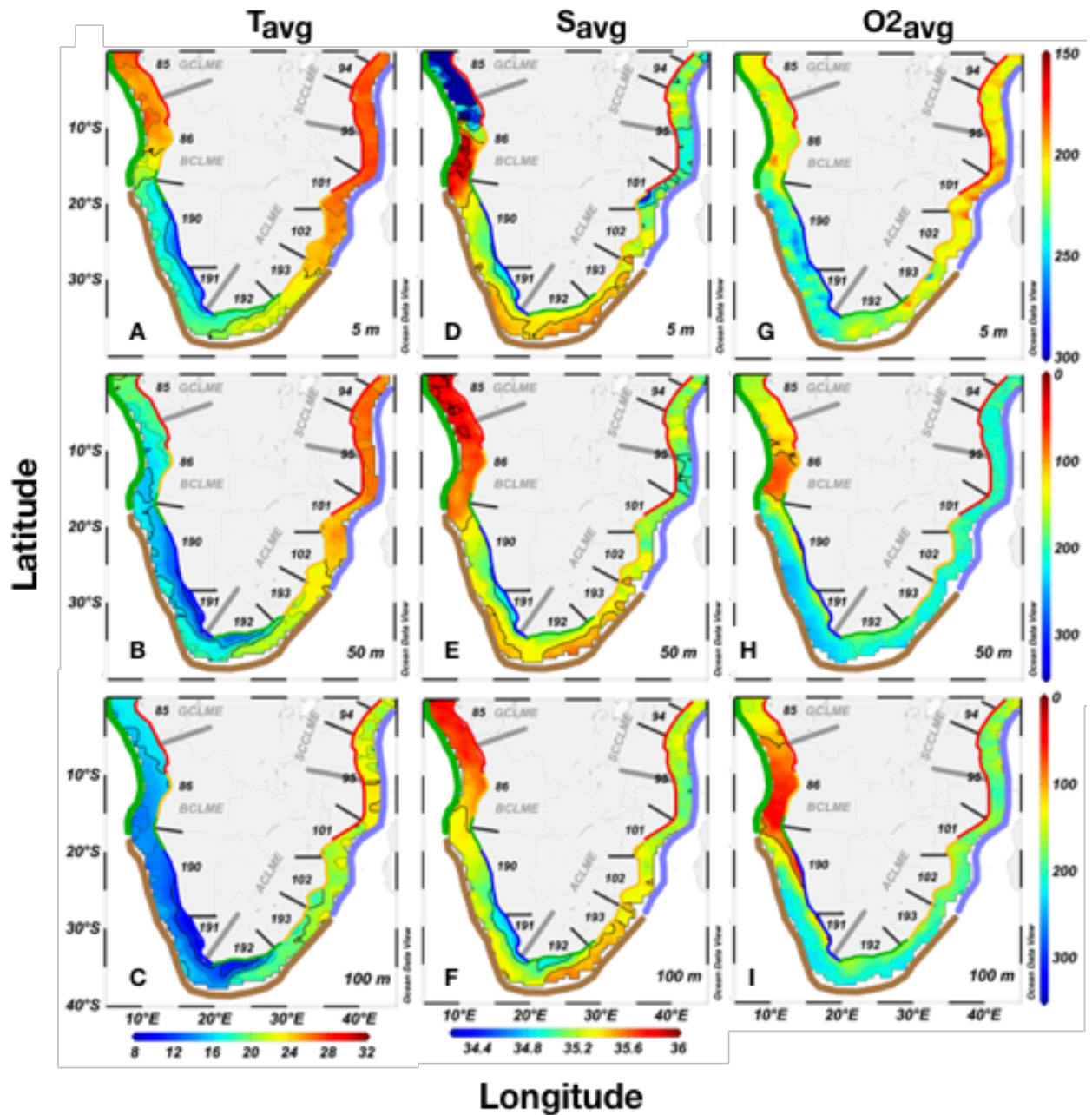
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Fig 8. . Seasonal oxygen climatologies at depths of 5 (A to D), 50 (E to H) and 100 m (I to L) for summer (JFM), autumn (AMJ), winter (JAS) and spring (OND). Contour intervals are 50  $\mu\text{M}$ , with 100  $\mu\text{M}$  indicated with a thicker line.

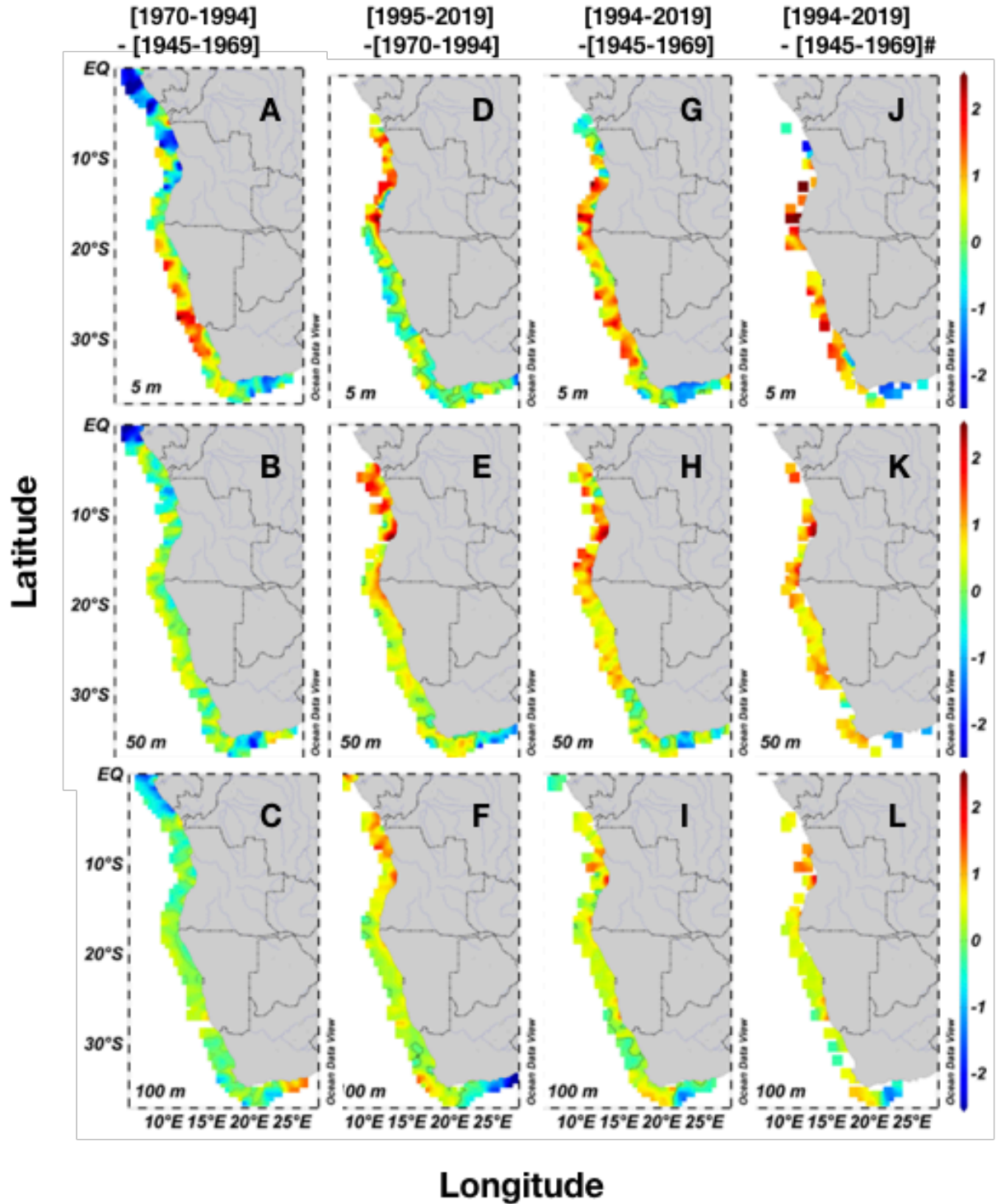
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Fig 9. Annual climatologies for temperature (A to C), salinity (D to F) and oxygen (G to I), at depths of 5, 50 and 100 m depths. Temperature contour intervals are 2°C, with 16, 20 and 24°C indicated with thicker lines. Salinity contour intervals are 0.2, with 34.6, 35.0 and 35.4 indicated with thicker lines. Oxygen contour intervals are 50  $\mu\text{M}$  and 100  $\mu\text{M}$  is indicated with a thicker line. The biogeographic boundaries of Fig 2 are superimposed on the climatologies.

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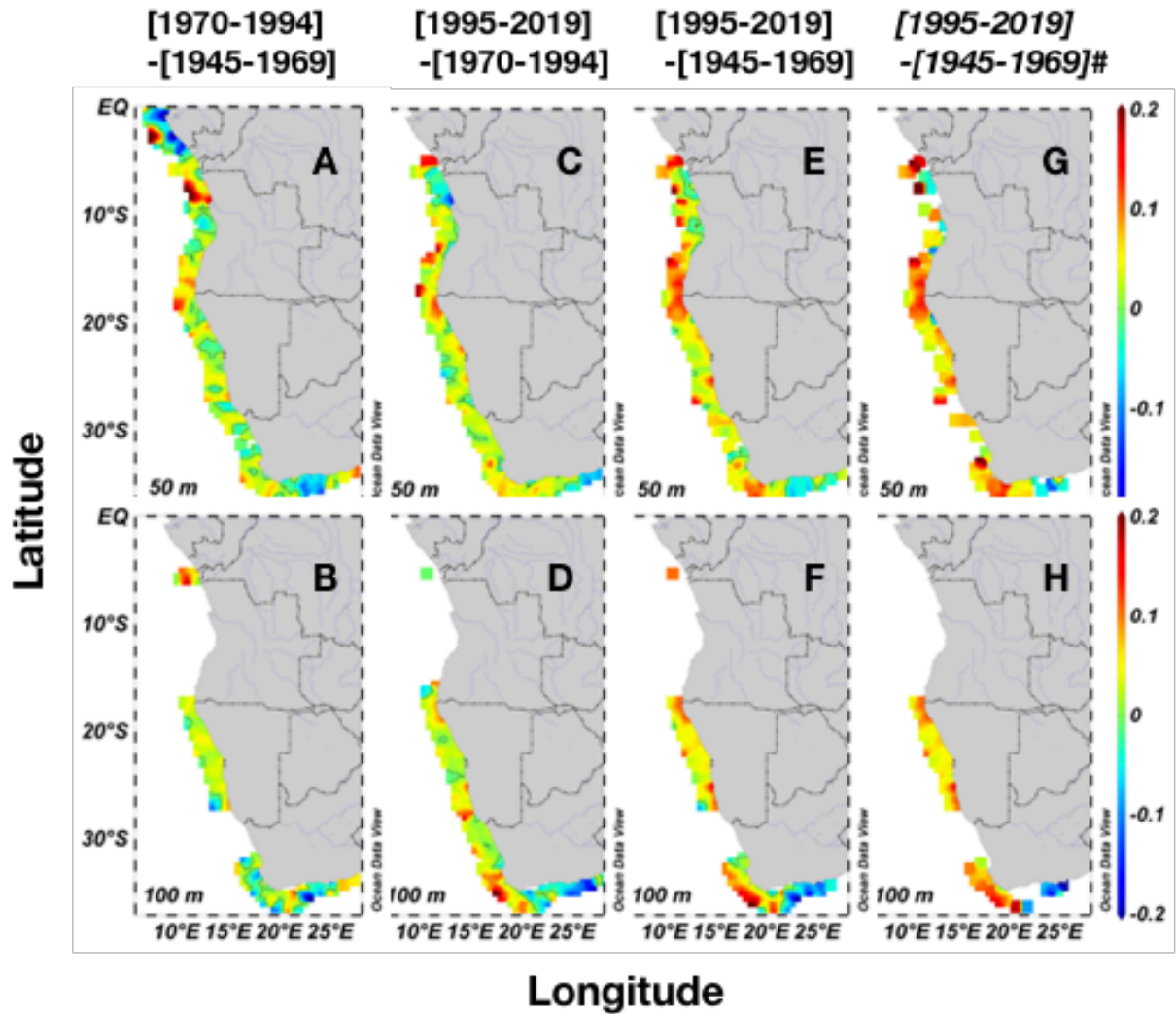


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Fig 10. The difference between 25 year climatological average annual temperature values ( $T_{25yr \text{ period}}$ ) at depths of 5, 50 and 100 m, for the following periods: A to C:  $T_{[1970 \text{ to } 1994]} - T_{[1945 \text{ to } 1969]}$ ; D to F:  $T_{[1995 \text{ to } 2014]} - T_{[1970 \text{ to } 1994]}$ ; G to I:  $T_{[1995 \text{ to } 2014]} - T_{[1945 \text{ to } 1969]}$ ; J to L:  $T_{[1995 \text{ to } 2014]} - T_{[1945 \text{ to } 1969]}^\#$ , with # indicating the exclusive plotting of H bins for which 25 year climatological values are available for each of the three 25 year periods, and that display monotonic trends (increasing or decreasing) from 1945 to 2014. The zero contour line is shown.



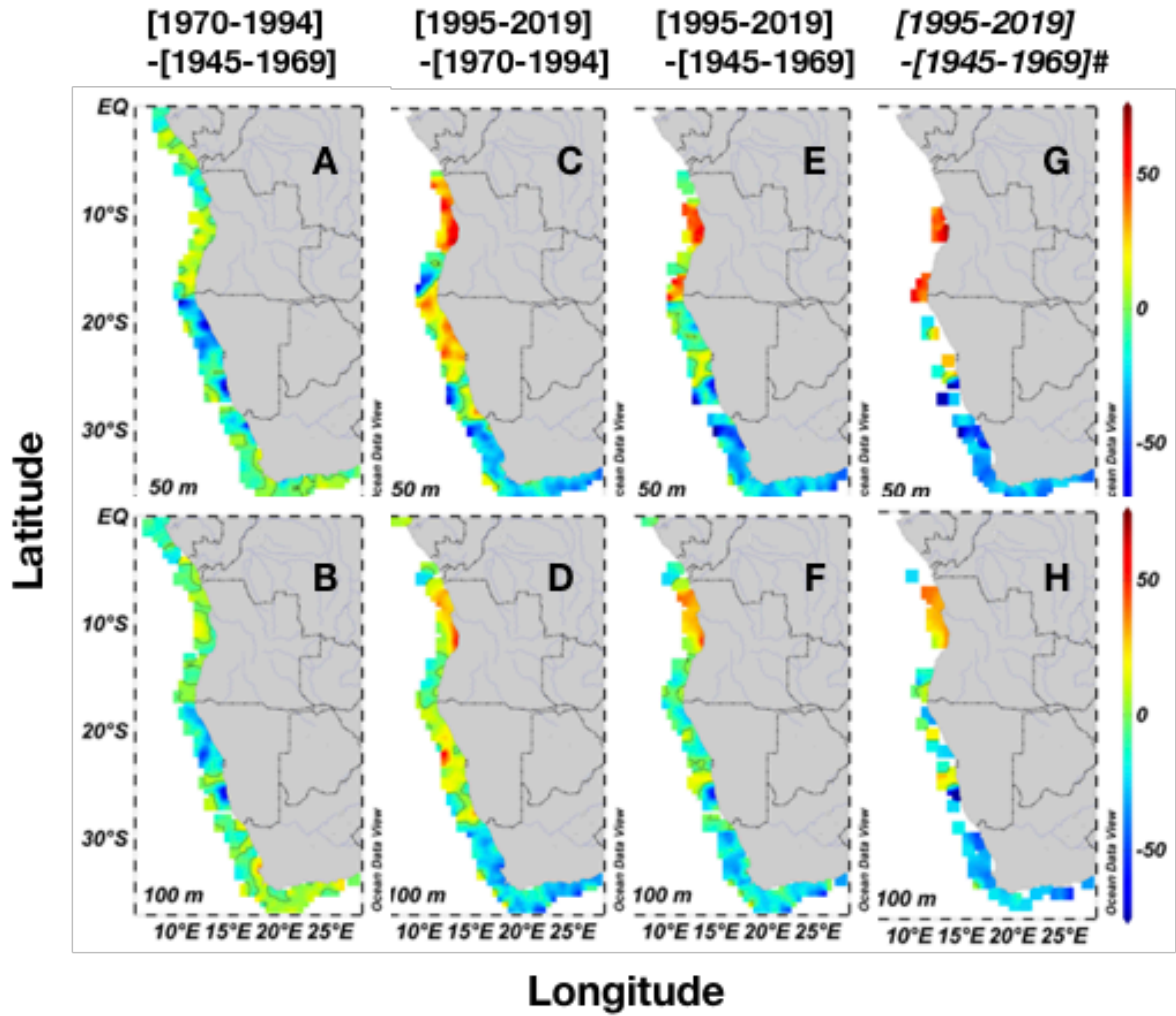
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Fig 11. The difference between 25 year climatological average annual salinity values ( $S_{25yr\ period}$ ) at depths of 5, 50 and 100 m, for the following periods: A to C:  $S_{[1970\ to\ 1994]} - S_{[1945\ to\ 1969]}$ ; D to F:  $S_{[1995\ to\ 2014]} - S_{[1970\ to\ 1994]}$ ; G to I:  $S_{[1995\ to\ 2014]} - S_{[1945\ to\ 1969]}$ ; J to L:  $S_{[1995\ to\ 2014]} - S_{[1945\ to\ 1969]}^\#$ , with # indicating the exclusive plotting of H bins for which 25 year climatological values are available for each of the three 25 year periods, and that display monotonic trends (increasing or decreasing) from 1945 to 2014. The zero contour line is shown.

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Fig 12. The difference between 25 year climatological average annual oxygen values ( $O_{25yr \text{ period}}$ ) at depths of 5, 50 and 100 m, for the following periods: A to C:  $O_{[1970 \text{ to } 1994]} - O_{[1945 \text{ to } 1969]}$ ; D to F:  $O_{[1995 \text{ to } 2014]} - O_{[1970 \text{ to } 1994]}$ ; G to I:  $O_{[1995 \text{ to } 2014]} - O_{[1945 \text{ to } 1969]}$ ; J to L:  $O_{[1995 \text{ to } 2014]} - O_{[1945 \text{ to } 1969]}^\#$ , with # indicating the exclusive plotting of H bins for which 25 year climatological values are available for each of the three 25 year periods, and that display monotonic trends (increasing or decreasing) from 1945 to 2014. The zero contour line is shown.

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**Appendix 2** WOD-18 data points excluded from the climatology, for each of the variables studied. The unique WOD Station ID can be used to find ancillary data for the particular station, at [ncei.noaa.gov](http://ncei.noaa.gov).

WOD StationID	Depth	Temperature	Salinity	Oxygen
18681	0	3,13		
167237	100		36,24	
178684	0		32,39	
355657	50			32
404741	0		32,30	
406622	0		33,06	
415965	0		33,49	
454521	0			26
465504	50			
532511	0		31,99	
598804	0	2,47	3,43	
717517	0			61
730626	0	7,44		
794626	0			
855259	0			
1323552	50		32,85	
6624740	0			676
6695714	0	3,02	2,69	
6737152	50	4,80		
6768833	0	2,64		
6794839	50	1,81		
6794839	100	1,81		
6796221	50	3,37		
6875195	0		37,78	
6887273	0		33,12	
7067809	0	35,01		
7213009	50		2,09	
7213009	100		1,80	
8337232	100		31,92	
8552259	0		33,35	
8552354	100			
8552370	100			
8552384	100			
8552597	0			
8552597	100			
8552597	0			
8552598	0			
8552691	0			
8552692	0			
8552693	0			
8552694	0			
8552696	0			
8552697	0			
8552698	0			
8552699	0			
8552748	0			
8552752	0			
8552754	0			
8552780	0			
8552794	0			
8552800	0			
8552994	0		34,01	
8553007	50			
8553010	50			
8553011	50			
8553012	50			
8553013	50			
8553014	50			
8553022	50			
8556262	0			
8556263	0			
8556264	0			
8556265	0			
8589195	100	4,06		
8589484	50			
8589485	50			
8589486	0			
8589486	50			
8589486	100			
8589487	50			
8589487	100			
8589488	50			
8589489	50			
8589490	50			
8589490	100			
8589491	50			
8589491	100			
8589558	0		30,59	
8589636	0	1,28		
8589671	50			
8589671	50			
8749324	50			1520
8749324	100			1517
8754127	50	0,90	0,90	
8754222	50		2,48	
8758038	100		38,32	
8758038	50		40,46	
9921458	0	2,48		
9921975	0	4,98		
10474743	0		15,09	
10474744	0		24,52	
10474745	0		27,99	
10474746	0		22,79	
10474747	0		11,58	
10474748	0		13,32	
10474749	0		22,06	
10474750	0		6,49	
10474751	0		24,98	
10474752	0		14,69	

**Appendix 3** Surface temperature 2-hour averages for all H bins with n > 200, for each of the four seasons, with the average of these 2-hr averages, T<sub>durnal</sub>, the average of all surface temperature values for that season, T<sub>all</sub>, and the difference between these averages. Also shown are the total number of casts for each season, and the distribution of these casts between the four daily 6-hour periods as shown. Temperature averages in brackets are bins with n < 3.

H type	Longitude	Latitude	Season	%casts				Tavg				Tavg				T <sub>durnal</sub>	T <sub>all</sub>	T <sub>all</sub> - T <sub>durnal</sub>						
				00:00-05:59	06:00-11:59	12:00-17:59	18:00-23:59	00:00-05:59	06:00-11:59	12:00-17:59	18:00-23:59	00:00-05:59	06:00-11:59	12:00-17:59	18:00-23:59									
Hcoast	11.75	-4.75	JFM	926	2	93	4	1	(27.25)	27.210	26.669	27.219	27.332	27.847	28.362	28.517	28.299	27.710	27.742	27.800	27.619	27.479	-0.140	
			AMJ	921	3	93	5	1	notdata	26.579	27.439	27.439	27.439	27.439	27.439	27.439	27.439	27.439	27.439	27.439	27.439	27.439	27.439	0.000
			JAS	888	1	93	4	1	20.900	(20.71)	20.528	21.217	21.240	21.543	22.512	22.329	22.407	22.210	21.747	(21.81)	21.747	21.391	21.547	-0.156
			OND	858	1	95	3	1	26.210	(25.45)	24.696	25.833	25.199	24.998	25.002	24.764	25.727	26.223	26.237	(26.02)	25.439	25.168	(25.07)	-0.271
			JFM	171	23	29	15	33	19.531	19.058	18.436	18.296	19.295	20.274	20.108	19.506	18.841	19.185	18.899	19.815	19.815	19.279	19.288	-0.009
H1*	11.25	-17.75	JFM	87	23	28	20	21	17.104	(16.82)	17.186	17.226	17.186	17.226	17.186	17.226	17.186	17.226	17.186	17.226	17.186	17.226	17.186	0.000
			AMJ	59	13	34	38	16	14.684	(14.45)	(14.27)	14.382	14.340	15.018	15.014	15.556	15.053	15.103	14.890	15.427	14.849	14.891	0.042	
			JAS	46	15	32	43	11	13.528	(14.03)	14.533	(13.920)	14.400	14.217	14.051	14.004	14.070	14.610	14.590	13.883	14.125	14.003	-0.122	
			OND	108	23	20	29	28	16.334	15.809	16.164	16.551	14.944	15.632	16.582	16.305	15.795	15.761	16.045	16.186	15.427	16.809	16.046	0.046
			JFM	111	18	10	37	35	19.461	18.863	17.806	(18.005)	18.146	19.361	19.248	17.491	17.752	17.302	18.041	18.928	18.367	18.183	-0.184	
H1*	12.75	-19.75	AMJ	82	18	38	18	26	16.753	16.189	15.508	(15.524)	15.886	16.041	16.655	15.470	15.627	15.370	16.122	15.805	15.763	-0.043		
			JAS	79	29	25	28	19	14.190	(14.46)	14.735	15.492	14.803	14.866	14.820	14.985	14.557	14.662	14.543	14.230	14.495	14.573	0.058	
			OND	63	28	25	22	26	14.130	13.867	13.722	(13.848)	13.974	14.058	14.960	14.963	14.735	14.672	15.747	14.366	14.420	14.512	0.092	
			JFM	125	18	23	32	27	18.427	18.651	18.988	(18.804)	19.841	20.421	19.403	19.763	19.546	18.383	19.001	19.377	19.170	19.297	0.127	
			AMJ	66	25	27	27	24	17.087	17.180	16.548	17.175	17.674	17.093	17.059	16.346	(16.705)	17.064	16.362	16.281	16.881	16.940	0.059	
Hcoast	14.25	-22.75	JFM	154	29	38	11	22	16.890	16.258	17.020	16.848	16.929	16.682	15.964	17.006	18.178	18.117	17.608	17.127	17.052	17.043	-0.004	
			AMJ	98	23	30	32	15	16.016	15.682	15.225	14.010	14.439	14.748	15.140	14.350	14.036	14.323	14.433	15.318	14.768	14.822	0.054	
			JAS	79	29	25	28	19	12.909	12.730	12.472	12.617	12.899	13.206	13.178	12.494	12.972	13.668	13.153	12.527	12.885	12.943	0.058	
			OND	43	16	47	23	14	14.137	(16.01)	15.910	14.439	14.140	14.380	14.014	(14.087)	14.160	13.943	13.727	(13.27)	14.193	14.181	-0.012	
			JFM	345	26	24	36	14	14.724	13.879	13.956	14.139	14.777	14.510	14.798	14.841	14.451	14.113	14.391	14.557	14.426	14.439	-0.017	
H1*	18.25	-32.25	AMJ	258	29	22	33	16	13.552	13.469	13.327	12.945	13.762	14.053	13.808	13.788	13.501	13.423	13.440	13.559	13.552	13.509	-0.047	
			JAS	226	28	37	27	9	13.107	12.898	12.994	13.054	13.284	13.521	13.469	13.286	13.256	13.600	12.743	13.100	13.177	13.077	0.101	
			OND	193	31	29	31	13	13.527	13.057	13.140	13.079	13.140	13.140	13.140	13.140	13.140	13.140	13.140	13.140	13.140	13.140	0.000	
			JFM	359	12	46	32	10	15.549	15.088	14.969	14.439	14.883	15.261	15.179	15.365	15.483	15.033	15.628	14.603	15.123	15.103	-0.020	
			AMJ	278	12	51	27	10	14.223	14.233	13.907	14.115	14.519	14.508	14.386	14.151	14.055	14.183	14.104	13.578	14.171	14.257	0.086	
H1*	17.75	-32.25	JAS	304	15	49	30	7	(13.48)	13.644	14.051	13.878	13.989	13.919	14.166	14.106	13.99	14.053	13.514	13.320	13.843	13.992	0.149	
			OND	251	20	43	25	14	14.631	14.089	13.923	13.923	14.089	14.141	14.759	15.030	14.499	14.289	14.607	14.710	14.549	14.549	-0.001	
			JFM	248	16	50	30	5	15.831	16.853	14.657	15.277	15.825	16.245	16.961	16.871	16.951	16.766	(17.27)	17.772	16.680	16.516	-0.164	
			AMJ	154	10	59	28	2	14.203	15.495	15.557	14.897	15.279	15.461	15.873	15.437	15.528	(15.200)	15.375	15.375	15.474	15.406	0.032	
			JAS	170	13	57	26	4	14.547	14.782	14.901	14.496	14.741	14.878	14.862	14.469	14.523	(15.540)	(14.85)	14.160	14.762	14.816	0.054	
H2*	16.75	-32.25	OND	151	12	60	25	6	16.645	15.605	15.605	15.728	15.625	15.417	15.711	15.508	15.767	(15.175)	16.065	16.637	16.745	16.745	-0.001	
			JFM	214	28	40	25	6	17.951	17.551	17.921	18.340	18.406	18.758	19.308	18.855	18.043	17.762	18.24	17.672	18.204	18.370	0.166	
			AMJ	127	21	30	43	7	(17.25)	17.444	16.884	16.884	17.045	16.572	17.275	17.87	16.903	17.270	16.983	17.072	17.121	17.072	-0.049	
			JAS	136	26	35	31	8	14.873	15.495	15.400	15.179	15.261	15.158	15.271	15.895	15.511	15.146	14.873	(14.87)	15.245	15.410	0.165	
			OND	110	25	24	45	6	(16.82)	16.530	16.190	16.082	16.425	16.578	16.837	16.728	15.812	16.380	(16.60)	16.478	16.574	0.096		
Hcoast	18.25	-32.75	JFM	271	17	31	16	54	13.477	13.113	13.167	13.167	13.167	13.167	13.167	13.167	13.167	13.167	13.167	13.167	13.167	13.167	0.000	
			AMJ	208	14	10	25	51	13.355	13.758	13.101	13.884	14.05	13.935	13.705	13.509	13.703	13.393	13.527	13.513	13.619	13.561	-0.058	
			JAS	305	11	9	28	52	13.020	12.988	(12.66)	12.414	13.298	13.672	14.346	13.386	13.386	13.386	13.386	13.386	13.386	13.386	13.386	0.000
			OND	235	15	11	22	53	13.365	12.996	(12.97)	12.94	13.906	14.469	14.268	14.336	13.198	13.413	13.428	13.403	13.558	13.561	0.004	
			JFM	292	22	28	34	16	14.479	14.312	14.154	14.041	14.599	14.670	14.727	14.548	14.551	14.371	14.002	14.953	14.434	14.490	0.058	
H1*	17.75	-32.75	AMJ	253	18	33	35	15	13.723	14.183	14.077	14.294	14.013	14.241	14.475	14.852	14.639	14.448	13.707	13.500	14.191	14.299	0.108	
			JAS	269	17	40	34	9	12.897	13.858	13.966	13.852	13.827	14.229	13.905	13.553	13.437	13.877	13.960	13.902	13.697	13.815	0.118	
			OND	246	23	29	32	16	12.872	13.807	13.671	13.606	14.155	14.660	14.117	13.990	14.344	13.991	14.329	14.016	13.913	13.990	0.077	
			JFM	254	20	37	33	10	17.755	17.21	16.949	17.346	17.111	17.062	17.223	17.157	16.887	17.091	17.951	17.100	17.235	17.276	0.041	
			AMJ	201	15	40	30	8	14.993	15.61	14.993	15.61	14.993	15.61	14.993	15.61	14.993	15.61	14.993	15.61	14.993	15.61	0.000	
H1*	17.5	-32.75	JAS	236	15	41	30	13	15.390	15.379	15.215	15.229	15.333	15.256	15.392	15.175	15.378	15.323	15.250	15.333	15.304	15.320	0.016	
			OND	217	17	37	31	15	16.084	16.265	16.472	16.388	16.593	16.716	16.758	16.849	16.849	16.849	16.849	16.849	16.849	16.849	16.849	0.000
			JFM	207	27	39	28	6	19.206	18.787	18.731	18.676	18.540	18.472	18.178	18.888	18.494	19.412	18.110	(18.66)	18.729	18.761	0.035	
			AMJ	158	28	36	30	6	17.047	16.915	17.070	16.878	16.676	17.247	17.290	16.978	17.229	16.965	16.756	17.153				

**Appendix 4** The standard deviation associated with the average of all available temperature data, and the standard deviation associated with the average of the the four seasonal climatological averages, at depths of 5, 50 and 100 m.

