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3 **Characterizing and Correcting Phase Biases in Short-Term, Multilooked**
4 **Interferograms**
5

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11 **Key Points: InSAR, Phase Bias, Fading Signal, Correction**

12 **Abstract**

13 Interferometric Synthetic Aperture Radar (InSAR) is widely used to measure deformation of the
14 Earth's surface over large areas and long time periods. A common strategy to overcome coherence
15 loss in long-term interferograms is to use multiple multilooked shorter interferograms, which can
16 cover the same time period but maintain coherence. However, it has recently been shown that
17 using this strategy can introduce a bias (also referred to as a "fading signal") in the interferometric
18 phase, particularly over vegetated areas. We isolate the signature of the phase bias by constructing
19 daisy chain sums of short-term interferograms covering identical 1-year time periods, but using
20 interferograms of different time spans. This shows that the shorter interferograms are more
21 affected by this phenomenon and the degree of the effect also depends on ground cover types. We,
22 propose a method for correcting the phase bias, based on the assumption that the bias in an
23 interferogram is linearly related to the sum of the bias in shorter interferograms spanning the same
24 time. We tested the algorithm over a study area in western Turkey by comparing average velocities
25 against results from a phase linking approach that has been shown to be almost insensitive to the
26 phase bias. Our corrected velocities agree well with those from phase linking approach. Our
27 approach can be applied to global compilations of short-term interferograms and offers the

28 possibility of accurate long-term velocities without a requirement for coherence in long-term
29 interferograms.

30 **1 Introduction**

31 Interferometric Synthetic Aperture Radar (InSAR) is a powerful tool for monitoring ground
32 deformation associated with earthquakes, volcanoes, landslides, and anthropogenic activities (e.g.
33 Biggs et al. 2009; Foughnia et al. 2019; Juncu et al. 2017; Massonnet et al. 1995; Temtime et al.
34 2018; Walters et al. 2011; Weiss et al. 2020). The accuracy of the estimated deformation is
35 traditionally thought to depend on uncorrected tropospheric and ionospheric delays, errors in phase
36 unwrapping, uncertainties in knowledge of satellite position, phase decorrelation due to changes
37 in scattering behavior between successive images, and system noise. Most of these error terms are
38 associated with individual epochs and cancel out when calculating the wrapped loop closure phase
39 at full spatial resolution, $\Delta\varphi$, defined for three epochs (i, j, k) as:

$$40 \quad \Delta\varphi_{i,k} = |\varphi_{i,k} - (\varphi_{i,j} + \varphi_{j,k})|_{2\pi} \quad (1)$$

41 where $\varphi_{i,j}$, for example, is the phase difference for a pixel in the interferogram between epochs i
42 and j, and $| \cdot |_{2\pi}$ indicates that the result is given modulo 2π (i.e. wrapped) (Michaelides et al.
43 2019; Zwieback et al. 2016).

44 For full-resolution processing, the wrapped loop closure phase must be zero. If multilooking, or
45 other forms of spatial filtering, is carried out as part of the processing, then $\Delta\varphi$ will not be precisely
46 equal to zero. This is because the filtering adds a term to each interferogram, the aim of which is
47 reduce the noise term, which does not cancel in the closure phase calculation. This is not an issue
48 for applications provided that the expected value of this term is zero. However, De Zan et al.

49 (2015) showed that the expected value of the filtering term is non-zero for some ground cover
50 types.

51 Nonzero closure phases are a product of the spatial filtering and are mainly associated with the
52 scattering and electrical properties of the ground surface (De Zan et al. 2015; Michaelides et al.
53 2019). Previous studies have suggested that changes in soil moisture and in the water content of
54 vegetation might lead to these phase inconsistencies (De Zan and Gomba 2018; De Zan et al.
55 2014). Although the amount of the bias caused by such inconsistencies is small in each individual
56 interferogram, its accumulation in time can significantly impact the final estimated velocities,
57 particularly for applications where millimetric accuracy is required. Ansari et al. (2021) showed
58 short-interval multilooked interferograms are more impacted by this phenomenon and referred to
59 this effect as the fading signal due to its short-lived nature. This is particularly problematic for
60 time-series analysis approaches that exploit pixels where coherence can only be maintained for
61 short time intervals – these pixels are likely to be strongly impacted by phase bias.

62 Mitigation strategies that have been proposed include correcting interferograms using physical
63 models such as a moisture-induced phase model (De Zan and Gomba 2018) or using “phase
64 linking” (PL) approaches, described below. Due to the varied sources of the phase bias,
65 employing a single physical moisture-induced phase model cannot account for all possible
66 sources of phase inconsistencies and no generic model exists to incorporate all possible sources
67 of the phase bias (Ansari et al. 2021). PL approaches, on the other hand, can effectively mitigate
68 this phenomenon by incorporating all possible $N(N - 1)/2$ interferometric phases obtained
69 from N SAR acquisitions (Guarnieri and Tebaldini 2008). The key step in the PL approaches is
70 to optimally estimate single-master phases for each pixel from all possible interferometric
71 combinations. These methods retrieve maximum available information in InSAR data stacks

72 (Samiei-Esfahany et al. 2016). Though efficient and robust, PL approaches require a large
73 number of interferograms and are computationally expensive, particularly for systems like
74 Sentinel-1, where there might be several hundred acquisitions. Moreover, the quality of the PL
75 estimated phases highly depends on the availability of the long-term interferograms. In case of
76 the decorrelated regions, the applicability and practicality of PL methods is limited.

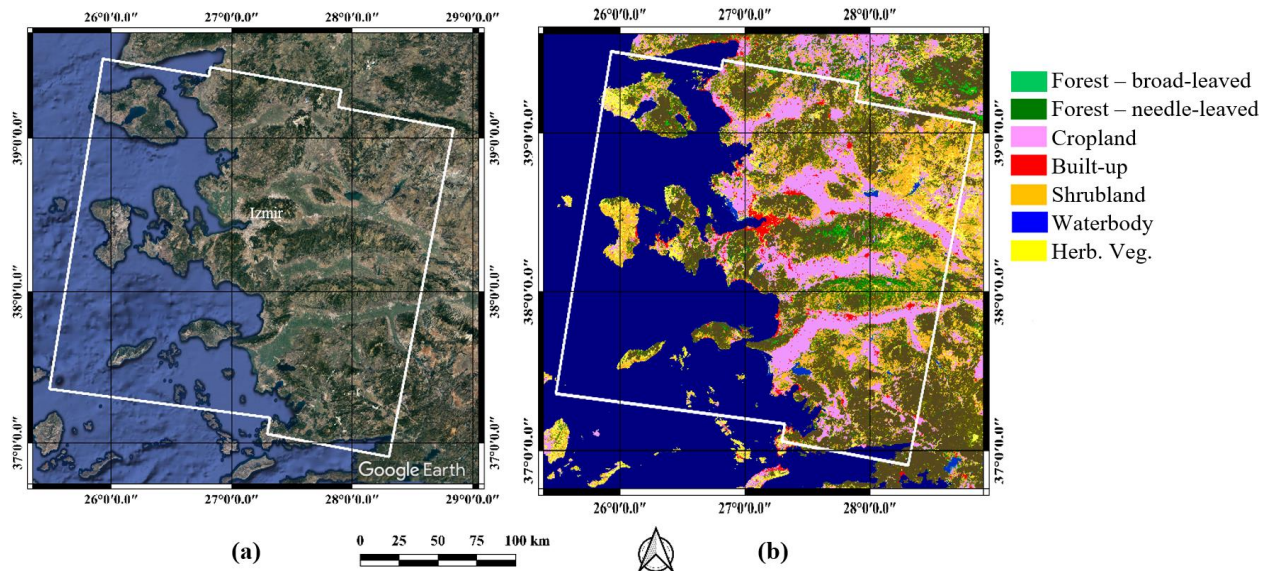
77 In this contribution, we first explore the characteristics of the phase bias by investigating its
78 temporal and spatial behavior. We then develop and test an empirical mitigation strategy to correct
79 short-term interferograms for the phase bias. Correcting for the phase bias in the short-term
80 interferograms is of great importance, in particular when the Small BAseline Subset (SBAS)
81 algorithms e.g. (Berardino et al. 2002; Morishita et al. 2020) are being used.

82 Our approach assumes that there is a linear relationship between the bias in a single interferogram
83 and the sum of the biases in the shorter interferograms spanning the same time. Employing this
84 assumption, we can estimate bias corrections for each interferogram through a linear least squares
85 inversion. We demonstrate the effectiveness of the proposed mitigation strategy by comparing the
86 resultant velocities with the phase linking approach.

87 **2 Study site**

88 We chose a study area in the west of Turkey that has a variety of ground cover types, including
89 forested and agricultural areas where long-term coherence is difficult to maintain (Figure 1).

90 Spatial heterogeneity in the land cover allows us to investigate the bias effect in these different
91 land covers ranging from more coherent urban areas to the agricultural and forest areas. The area
92 is imaged by Sentinel-1 A and B data on every overpass. We processed all interferometric pairs



93

94 Figure 1) Study site: (a) Overview of the study area located in the western Turkey. Izmir is the major city, situated along
 95 the Aegean coast. The white polygon shows the footprint of the Sentinel-1 data from descending path 36. (b) Land cover map
 96 obtained through the Copernicus Land Monitoring Service (<https://lcviewer.vito.be>).

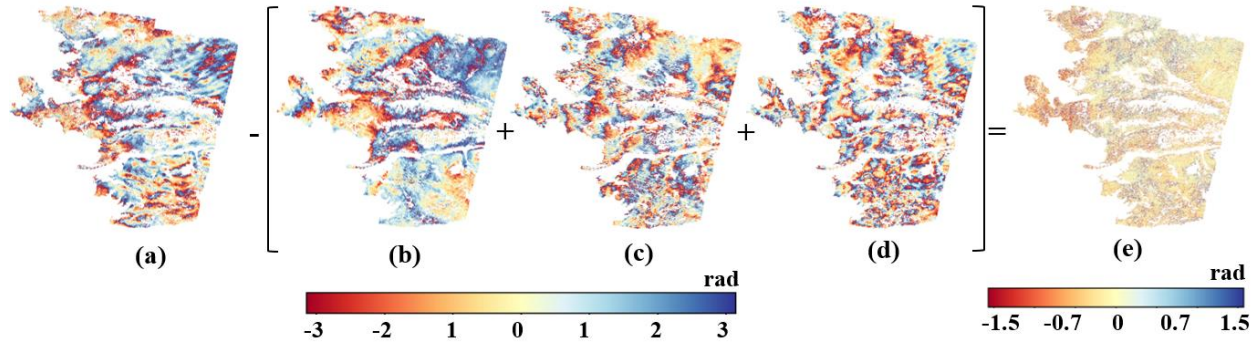
97 from one-year of Sentinel-1 acquisitions on track 36, where 60 images were acquired in the period
 98 spanning 1 February 2017 to 31 January 2018. All interferograms were generated using the
 99 automated workflows from the COMET-LiCSAR system (Lazecký et al. 2020), and were
 100 multilooked by factors of 5 in the range and 20 in the azimuth directions and geocoded onto a 100
 101 m grid using elevation data from the Shuttle Radar Topography Mission (Farr et al. 2007).

102 **3 Phase bias characterization**

103 Although the bias in individual interferograms cannot be isolated, we can measure phase bias in
 104 sets of interferograms by examining loop closure phases using different combinations of data.

105 Figure 2 shows how we calculate the closure phase using a set of multilooked interferograms in a
 106 loop. In this example, we subtracted the sum of three 6-day interferograms (b), (c) and (d) from
 107 an 18-day interferogram (a) to isolate the loop closure phase (e). We use the notation

108 $\Delta\varphi^{18-6} = 18 \text{ day} - \sum_1^3(6 \text{ day})$ to denote this loop closure phase. For the rest of this paper



109

110 Figure 2) Example of closure phase calculated from an 18-day interferogram and three 6-day interferograms. 18-day
 111 interferogram (a) spans 2017-02-18 to 2017-03-08. Three 6-day interferograms (b,c and d) span 2017-02-18 to 2017-02-24 (b),
 112 2017-02-24 to 2017-03-02 (c) and 2017-03-02 2017-03-08 (d). The resulting closure phase is shown in (e).

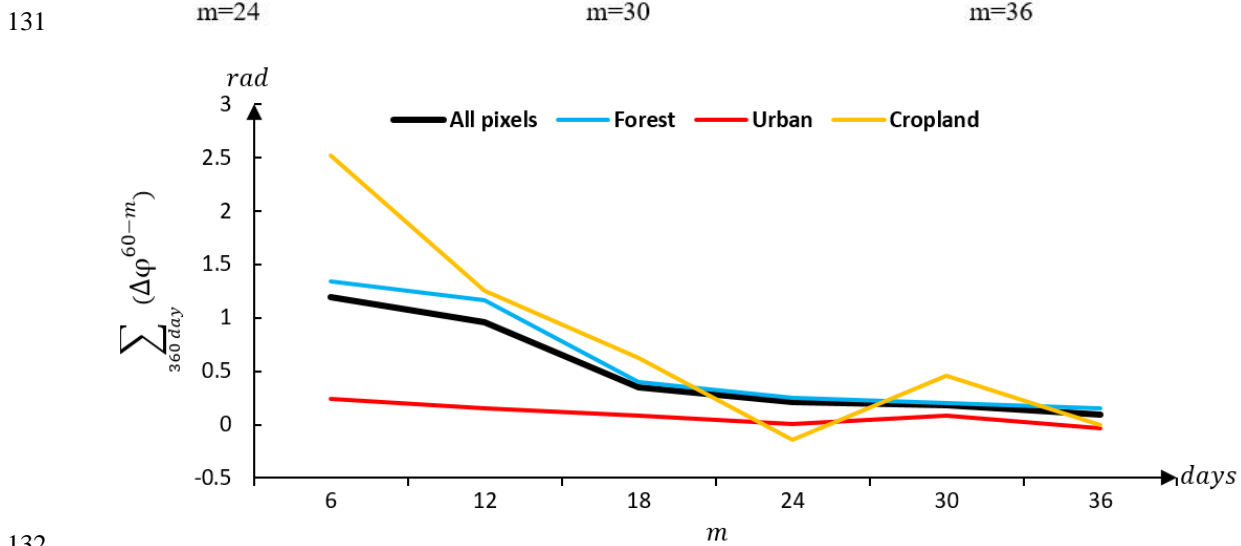
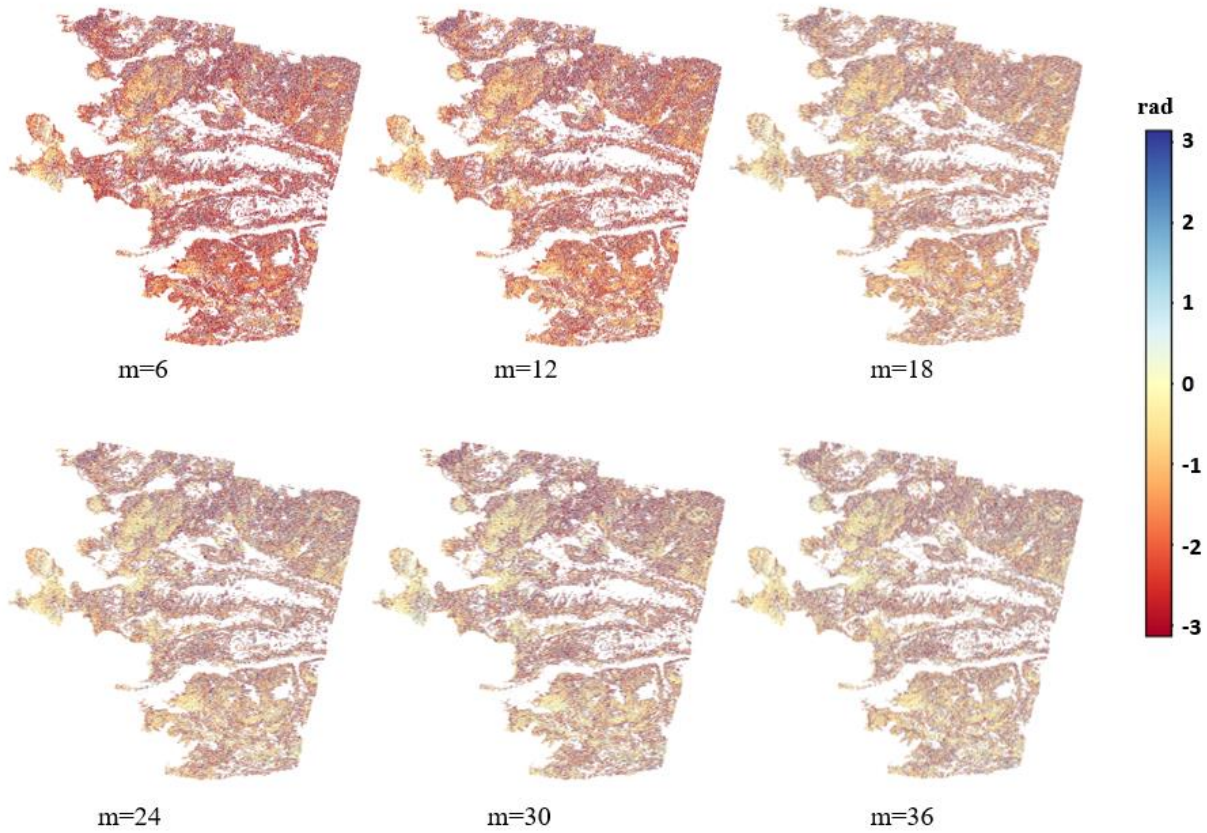
113 $\Delta\varphi^{n-m}$ indicates the loop closure phase from subtracting the summation of all m day
 114 interferograms from a n day interferogram spanning the same time. We also used the notation
 115 $\sum_{360 \text{ days}} \Delta\varphi^{n-m}$, for example, to show the 360-day cumulative loop closure phase calculated
 116 as the difference between 360-day “daisy chain” sums of interferograms with length n days and
 117 m days.

118 We use wrapped phases throughout this study to calculate the closure phases, with the result of
 119 any phase differences rewrapped to the interval $\pm\pi$. Taking a closer look at Figure 2(e), we can
 120 see a spatially correlated signal that varies across the image. Comparing this Figure with Figure
 121 1(b), this phase bias signal appears strongest in the vegetated areas.

122 To understand how the phase bias varies in interferograms of different lengths, we calculated the
 123 360-day cumulative loop closure phase using $n=60$ and $m=6, 12, 18, 24, 30$ and 36 . The results
 124 are shown in Figure 3(top).

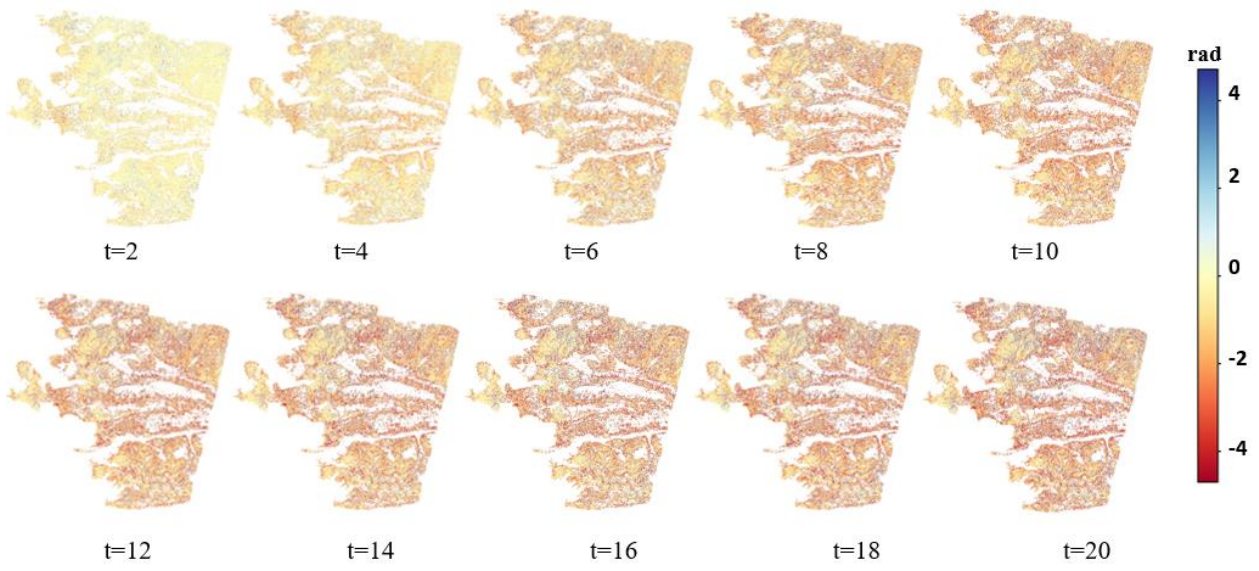
125 The results show that shorter interferograms are more affected by this phenomenon, with
 126 cumulative loop closure phases reducing in size dramatically as the length of the shorter
 127 interferograms in the loop increases. This observation agrees with the effect of the fading signal

128 (Ansari et al. 2021). The magnitude of the bias averaged over multiple pixels strongly depends
 129 on the ground cover type, with cropland and forested pixels having significantly larger bias than
 130 urban pixels (Figure 3 (bottom)).

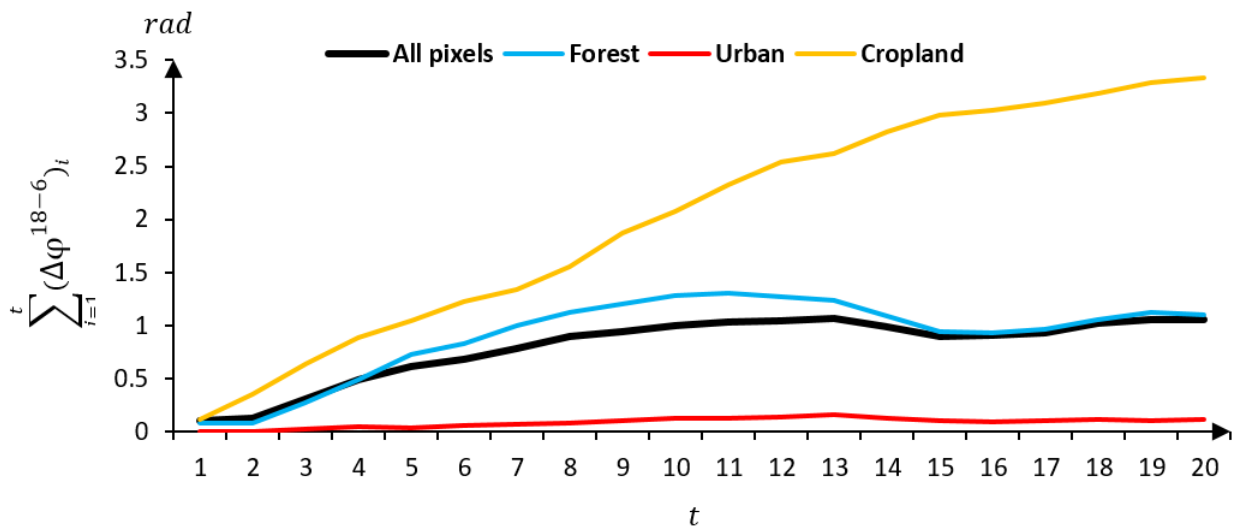


132
 133 Figure 3) 360-day cumulative loop closure phases, $\sum_{360 \text{ day}} (\Delta\varphi^{60-m})$, for varying timespans, m (top), and mean value of
 134 cumulative loop closure phases for different land cover classes as a function of m (bottom)

135 To test how the phase bias accumulates in time, we calculated $\sum_{i=1}^t (\Delta\varphi^{18-6})_i$ for $t=1, \dots, 20$, 20
 136 being the total number of consecutive 18 day interferograms in the 360-day observation period
 137 (Figure 4). The results show that although the amount of the closure phase is small in each
 138 individual loop, it increases with time. The rate of bias accumulation is not steady throughout the
 139 year, being highest for cropland and lowest for urban pixels.



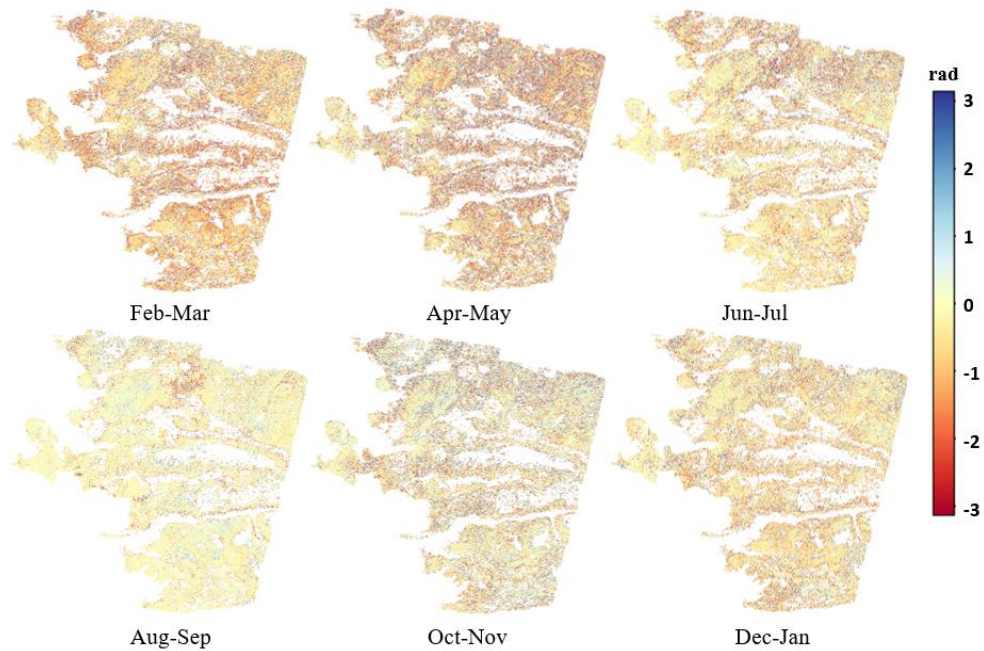
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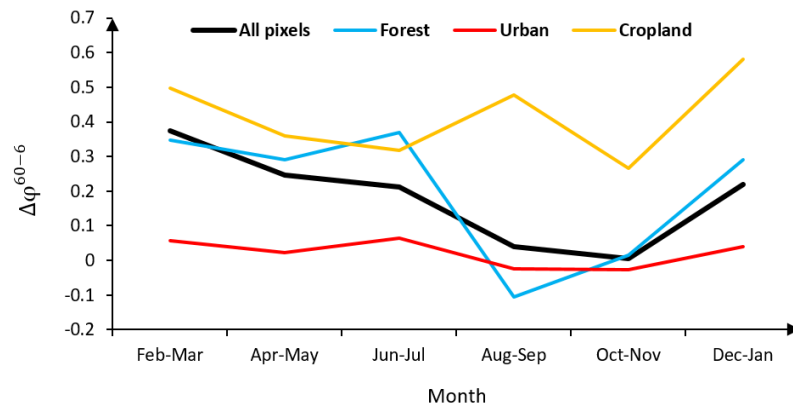
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142 Figure 4) The temporal accumulation of the loop closures obtained by $\sum_{i=1}^t (\Delta\varphi^{18-6})_i$ (top), and temporal accumulation of phase
 143 bias averaged for different land covers within the scene (bottom).

144 We also investigated the temporal variation in bias accumulation by calculating $\Delta\varphi^{60-6}$ loop
 145 closure phases covering different two-month periods (Figure 5). Each plot in Figure 5 (top)
 146 belongs to a two-month period. Figure 5 (bottom) illustrates the mean value of plots in different
 147 periods and in different landcovers. The plots indicate that the strength of the bias varies
 148 throughout the year. The largest mean value of phase bias is observed in the first plot, which
 149 corresponds to the period February and March. The smallest mean value, on the other hand,
 150 occur in late summer (August to September).

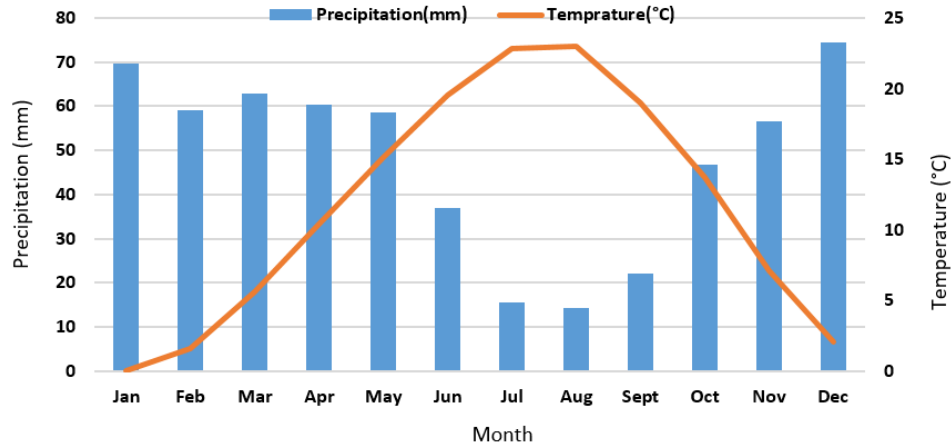


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152

153 Figure 5) Seasonal variation of the bias. The temporal plots of $\Delta\varphi^{60-6}$ for 1 year.



154

155 Figure 6) Average monthly temperature and rainfall of Turkey for 1991-2020 obtained from

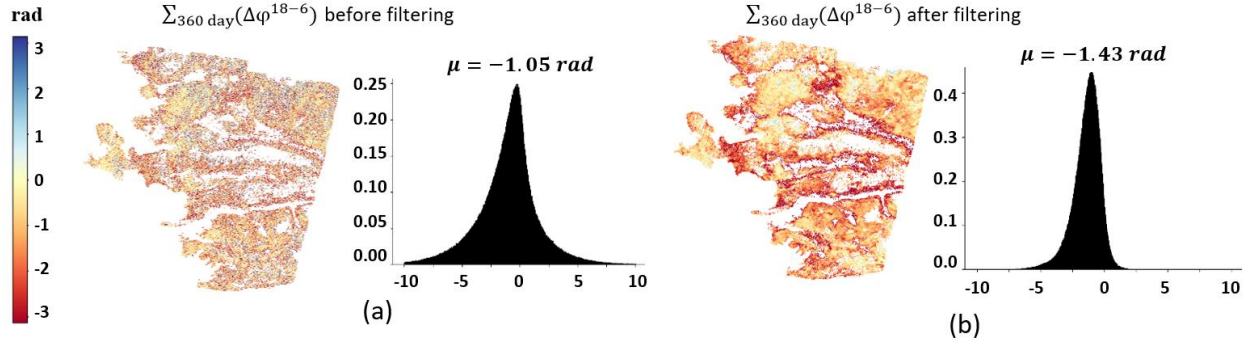
156 <https://climateknowledgeportal.worldbank.org/country/turkey/climate-data-historical>

157 This matches well with the precipitation season in the west coast of Turkey (highest in January to
 158 March and lowest in July to September) as shown in Figure 6. The closure phase in cropland
 159 pixels is more complex and may depend on several factors, including the vegetation growth as
 160 well as moisture variation.

161 Finally, in the last experiment of this section, we investigated the effect of the adaptive phase
 162 filtering (Goldstein and Werner 1998) on phase bias. Phase filtering is commonly applied to
 163 interferograms to reduce phase noise which greatly improves phase unwrapping performance. We
 164 applied a spatial filter to the multilooked interferograms using an adaptive power spectrum filter
 165 with FFT window size=32 and alpha=1. Figure 7 compares the cumulative loop closure phase
 166 $\sum_{360 \text{ day}}(\Delta\varphi^{18-6})$ using unfiltered and filtered interferograms. Filtering increases the mean value
 167 of the loop closure phase (the bias), by effectively increasing the multilooking factor. Therefore,
 168 we recommend caution in using filtered interferograms for time-series analysis.

169

170



171

172 Figure 7) $\Sigma_{360 \text{ day}}(\Delta\varphi^{18-6})$ and its histogram applied to the multilooked interferograms that are not filtered (a) and those that are
 173 (b). Here we used the adaptive Goldstein filtering.

174 4 Phase bias correction

175 In the COMET-LiCSAR automatic processing system (Lazecký et al. 2020), interferograms have
 176 been processed that connect each epoch, i , to the three or four nearest acquisitions in time,
 177 backward and forward. We therefore aim to develop a bias correction approach that uses just the
 178 interferograms formed from the closest three connections so that accurate velocities can be
 179 obtained without requiring mass processing of large numbers of additional longer-term
 180 interferograms.

181 Several loop closure phases can be calculated for an individual pixel, from these interferograms,
 182 including:

$$\Delta\varphi_{i,i+2} = \varphi_{i,i+2} - (\varphi_{i,i+1} + \varphi_{i+1,i+2}) \text{ and} \quad (2)$$

$$\Delta\varphi_{i,i+3} = \varphi_{i,i+3} - (\varphi_{i,i+1} + \varphi_{i+1,i+2} + \varphi_{i+2,i+3}), \quad (3)$$

183

184 where $\Delta\varphi_{i,i+2}$ and $\Delta\varphi_{i,i+3}$ are the $\Delta\varphi^{12-6}$ and $\Delta\varphi^{18-6}$ loop closure phases respectively. Assuming
 185 the closure phase is due to biases and noise in each interferogram, Equation (2) and (3) can be
 186 written as:

$$\Delta\varphi_{i,i+2} = \delta_{i,i+2} - (\delta_{i,i+1} + \delta_{i+1,i+2}) + \varepsilon \text{ and} \quad (4)$$

$$\Delta\varphi_{i,i+3} = \delta_{i,i+3} - (\delta_{i,i+1} + \delta_{i+1,i+2} + \delta_{i+2,i+3}) + \varepsilon, \quad (5)$$

187 where $\delta_{i,j}$ is the unknown phase bias in the interferogram formed from images i and j , and ε is the
188 sum of the noise terms.

189 If we want to solve for the unknown phase bias terms $\delta_{i,j}$ on each 6-, 12- and 18-day interferogram,
190 using the two sets of loop closure observations, $\Delta\varphi_{i,i+2}$ and $\Delta\varphi_{i,i+3}$, then with N acquisitions we
191 have $2N-5$ observations and $3N-6$ unknowns. The system of equations is therefore
192 underdetermined.

193 To solve this underdetermined inverse problem we introduce an assumption that the bias in an
194 interferogram is linearly related to sum of biases in shorter interferograms spanning the same time.
195 In other words, although the bias varies in strength with time, we assume the change in strength of
196 the bias in interferograms of different length is a constant ratio. i.e.

$$\delta_{i,i+2} = a_1 (\delta_{i,i+1} + \delta_{i+1,i+2}) \text{ and} \quad (6)$$

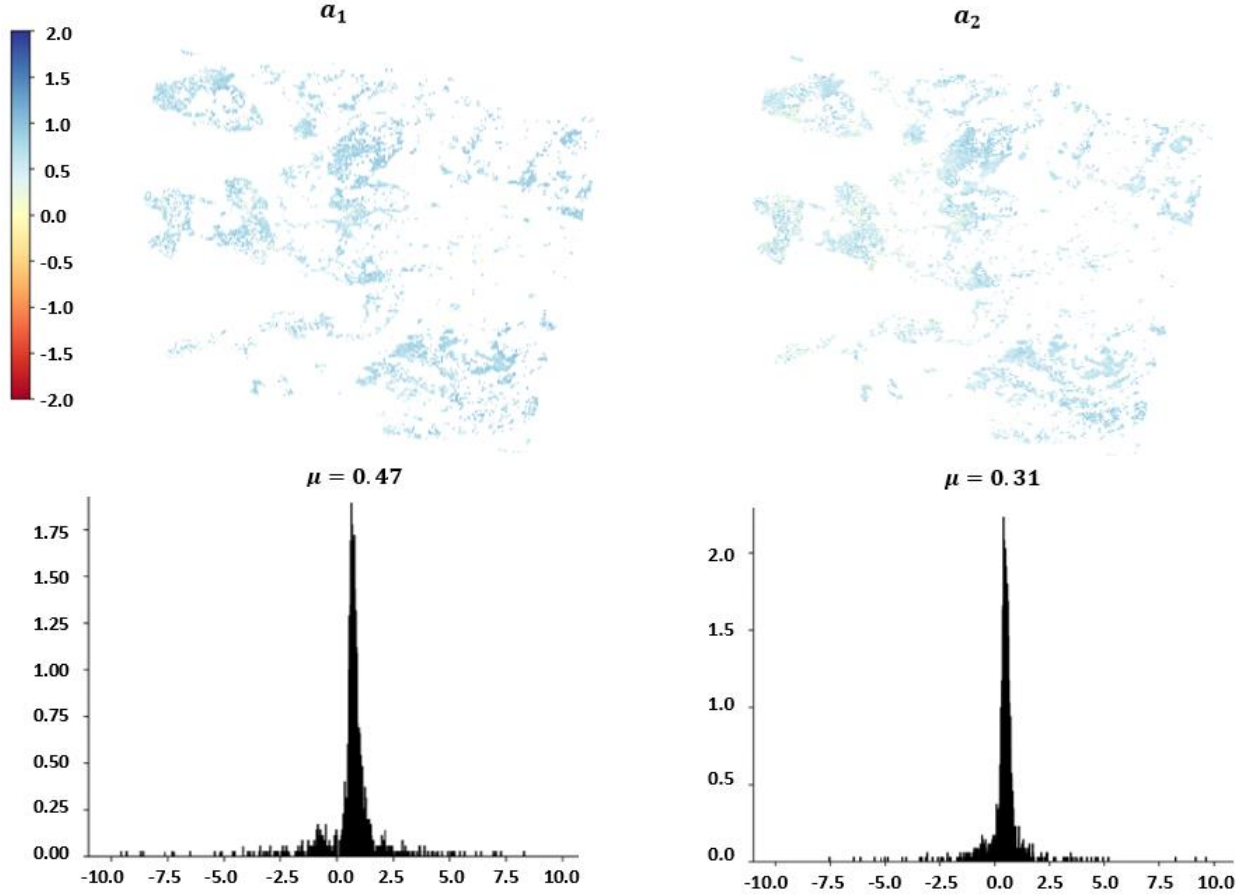
$$\delta_{i,i+3} = a_2 (\delta_{i,i+1} + \delta_{i+1,i+2} + \delta_{i+2,i+3}), \quad (7)$$

197 where a_1 and a_2 are unknown constants that linearly relate the bias in the longer interferograms
198 to the sum of the corresponding biases in the short interferograms covering the same time period.

199 If we assume that 360-day interferograms have negligible bias, a_1 and a_2 can be estimated for
200 each pixel by calculating the ratio of the cumulative loop closure phases for 12- and 6-day
201 interferograms and 18- and 6-day interferograms respectively:

$$a_1 = \frac{\Delta\varphi^{360-12}}{\Delta\varphi^{360-6}} \text{ and} \quad (8)$$

$$a_2 = \frac{\Delta\varphi^{360-18}}{\Delta\varphi^{360-6}} \quad (9)$$



202

203 Figure 8. The maps of a_1 and a_2 (top), and their corresponding histograms (bottom)

204 When estimating the regularization parameters a_1 and a_2 only pixels that remain coherent for a
 205 period of 1 year can be used. Figure 8 shows the maps of the a_1 and a_2 and their histograms.
 206 Although estimates of a_1 and a_2 for each pixel are noisy, there is no systematic pattern in space
 207 (Figure 8 (top)), suggesting that a single value is appropriate. In this study, we used the mean
 208 values of 0.47 and 0.31 for a_1 and a_2 respectively for all pixels.

209 If a_1 and a_2 are constants, using equations (4) to (7) and including all observations in epochs i to
 210 $i+3$ lead to a series of observation equations relating the closure phases to unknowns $\delta_{i,i+1}$.

$$\begin{pmatrix} \Delta\varphi_{i,i+2} \\ \Delta\varphi_{i+1,i+3} \\ \Delta\varphi_{i,i+3} \end{pmatrix} \cong \begin{pmatrix} a_1 - 1 & a_1 - 1 & 0 \\ 0 & a_1 - 1 & a_1 - 1 \\ a_2 - 1 & a_2 - 1 & a_2 - 1 \end{pmatrix} \begin{pmatrix} \delta_{i,i+1} \\ \delta_{i+1,i+2} \\ \delta_{i+2,i+3} \end{pmatrix}. \quad (10)$$

211 This reduces the number of unknowns to N-1, the biases for the 6-day interferograms. The system
 212 of equations is then overdetermined when $N \geq 5$ (with $2N-5$ loop closure observations) and the
 213 unknown bias terms can be estimated using a linear least squares inversion.

214 Upon the estimation of the bias terms, every single 6-, 12- and 18-day interferogram can then be
 215 corrected using

$$\varphi_{i,i+1}^c = \varphi_{i,i+1} - \hat{\delta}_{i,i+1} \text{ and} \quad (11)$$

$$\varphi_{i,i+2}^c = \varphi_{i,i+2} - \hat{\delta}_{i,i+2} = \varphi_{i,i+2} - a_1(\hat{\delta}_{i,i+1} + \hat{\delta}_{i+1,i+2}) \text{ and} \quad (12)$$

$$\varphi_{i,i+3}^c = \varphi_{i,i+3} - \hat{\delta}_{i,i+3} = \varphi_{i,i+3} - a_2(\hat{\delta}_{i,i+1} + \hat{\delta}_{i+1,i+2} + \hat{\delta}_{i+2,i+3}), \quad (13)$$

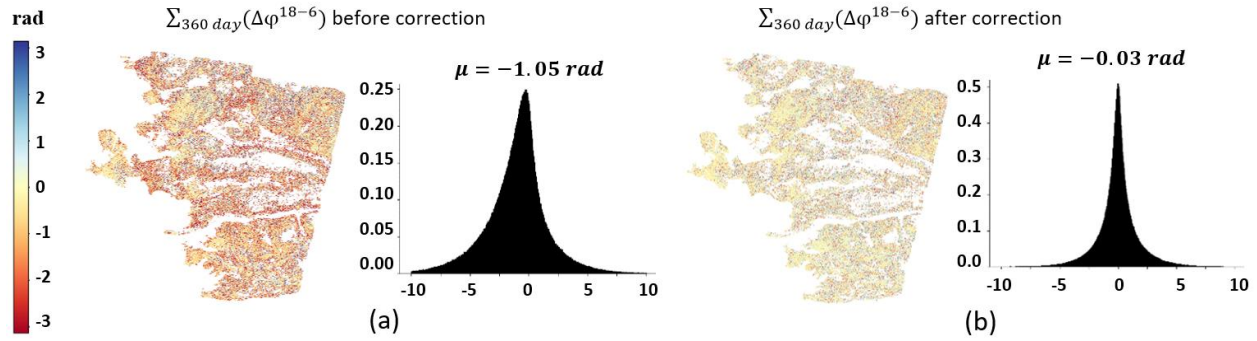
216 where $\varphi_{i,i+1}$, $\varphi_{i,i+2}$ and $\varphi_{i,i+3}$ are the original 6-day, 12-day and 18-day interferograms and $\varphi_{i,i+1}^c$,
 217 $\varphi_{i,i+2}^c$ and $\varphi_{i,i+3}^c$ are the corrected interferograms. The $\hat{\delta}_{i,j}$ are the estimated bias terms.

218 **5 Correction results**

219 All the experiments in this section were carried out on a set of coherent pixels, which were selected
 220 by applying a threshold of 0.3 on the 18-day average coherence. We estimated the corrections
 221 using Equations (10) and corrected all the 6-day 12-day and 18-day interferograms covering our
 222 360-day study period using Equations (11), (12) and (13) respectively.

223 Figure 9 shows a comparison between the closure phase $\sum_{360 \text{ day}}(\Delta\varphi^{18-6})$ using the original
 224 interferograms and that found using the corrected interferograms. It is clear that correcting the
 225 interferograms has significantly decreased the closure phase, with its mean and the standard
 226 deviation decreasing from $-1.05 \pm 2.7 \text{ rad}$ prior to correction to $0.03 \pm 1.7 \text{ rad}$.

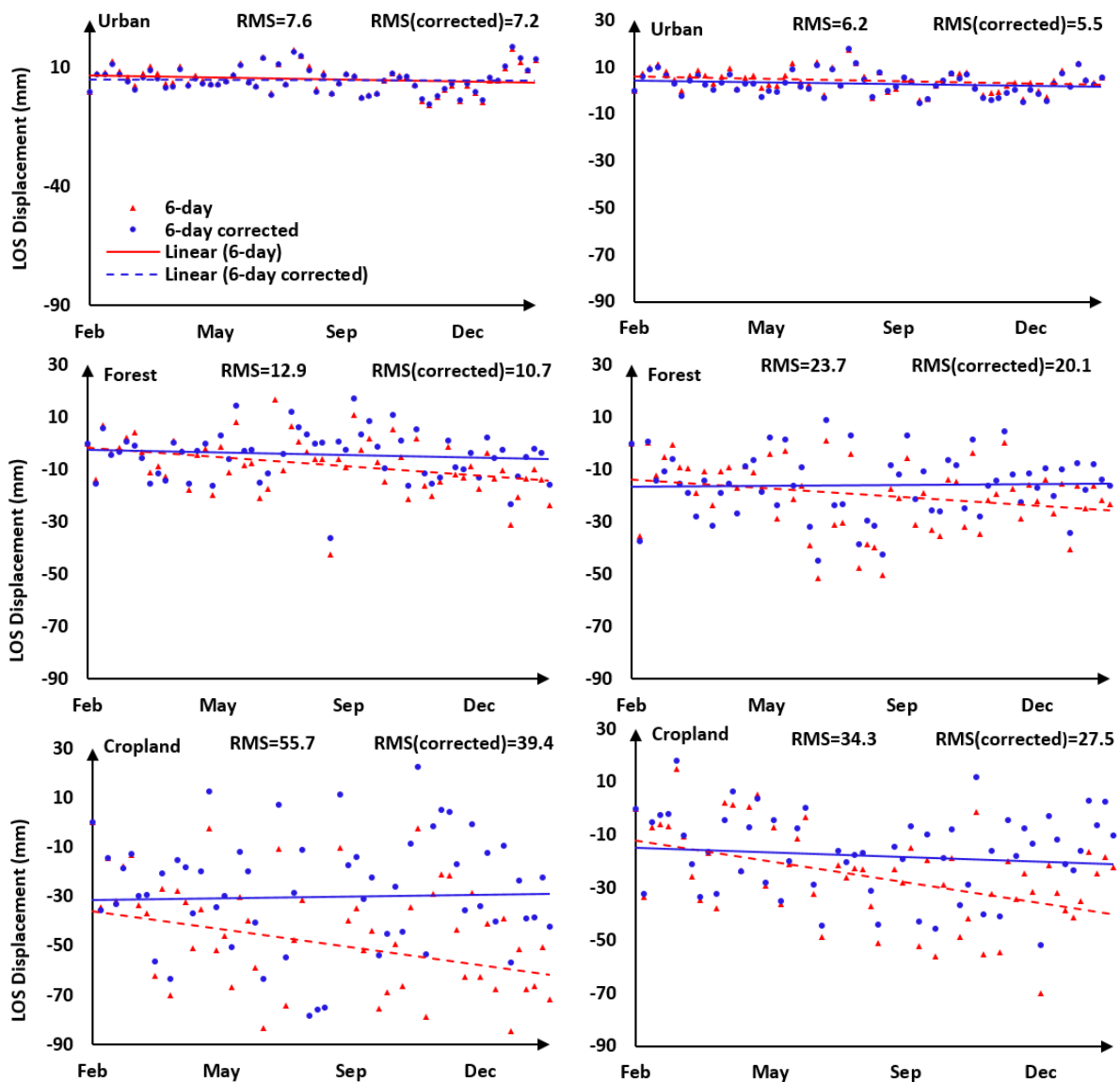
227 We also show (Figure 10) corrected and uncorrected time series of line of sight (LOS)
 228 displacement calculated from just the 6-day interferograms for some example points in different



229

230 Figure 9) The cumulative loop closure phase $\sum_{1yr}(18\text{ day} - 6\text{ day})$ calculated using (a) the original interferograms (b)

231 interferograms corrected with our empirical correction.



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235 Figure 10) The LOS time-series displacement for two example points in each land cover. RMS and RMS(corrected) show the
 236 root-mean-square of residuals calculated before and after correction respectively

237 land covers. Pixels in urban areas change the least with the correction, whereas the agricultural
238 pixels have larger values of corrections. We also calculate the root-mean-square (RMS) of the
239 residuals before and after the correction (Figure 10); the correction reduces the scatter of the data
240 for all land cover types. Considering all the pixels, the mean RMS residual of fit to the linear time
241 series model has reduced from 27.2 rad before the bias correction to 20.7 rad after correcting for
242 the phase bias.

243 **6 Validation**

244 As well as demonstrating the reduction in cumulative loop closure phases, we can also compare
245 line-of-sight velocities estimated from our corrected and uncorrected data with velocities from an
246 approach that is less sensitive to phase bias. We use a Phase Linking (PL) approach for this
247 validation test, which uses all possible interferograms and has been shown to be rather unaffected
248 by phase bias of short-term interferograms (Ansari et al. 2021). There are a number of PL methods
249 in the literature. These methods try to obtain the best estimates of $N - 1$ phase differences for a
250 pixel relative to the primary date using the $N(N - 1)/2$ available interferometric phases. PL
251 methods are categorized into maximum-likelihood estimators (Ferretti et al. 2011), least squares
252 estimators (Samiei-Esfahany et al. 2016), Eigen decomposition-based phase estimators (Cao et al.
253 2016; Fornaro et al. 2015) and Eigen decomposition-based Maximum-likelihood estimator (Ansari
254 et al. 2018).

255 Eigen decomposition-based methods are relatively computationally efficient and straightforward
256 to implement; we use the approach from (De Zan et al. 2007; Fornaro et al. 2015), hereafter

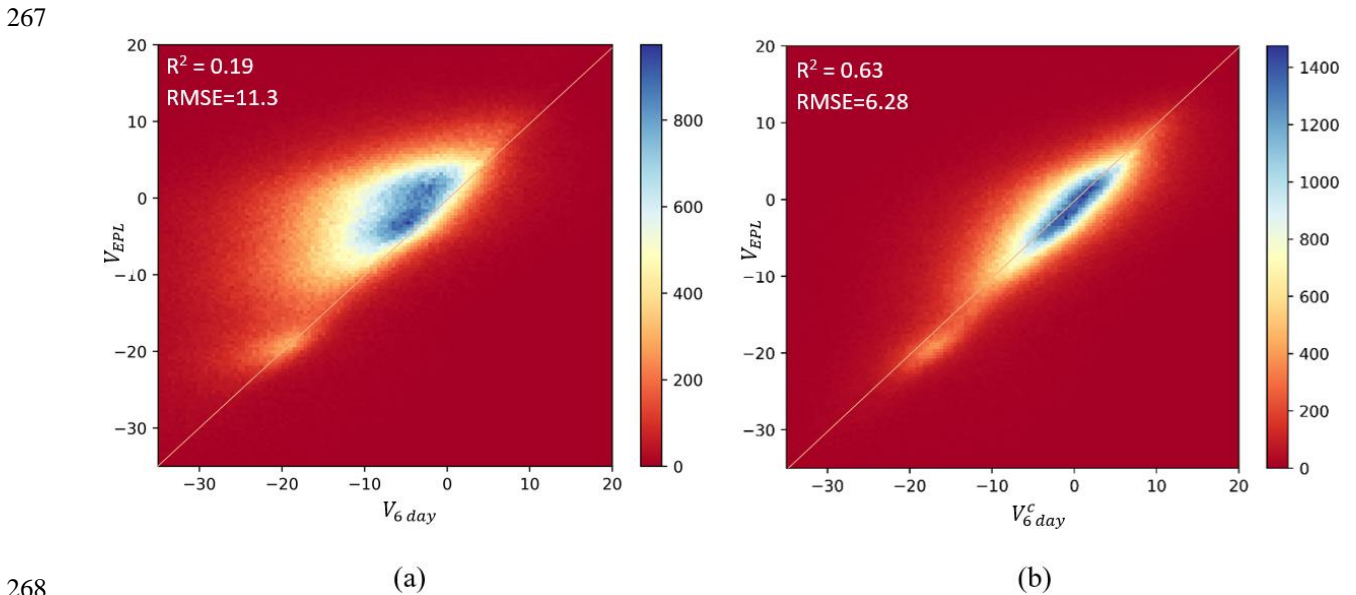
257 referred to as EPL, as our reference method to compare with results from our inversion that only
 258 uses short-interval interferograms.

259 Considering T as the N by N coherence matrix, the Eigen decomposition of T can be obtained as

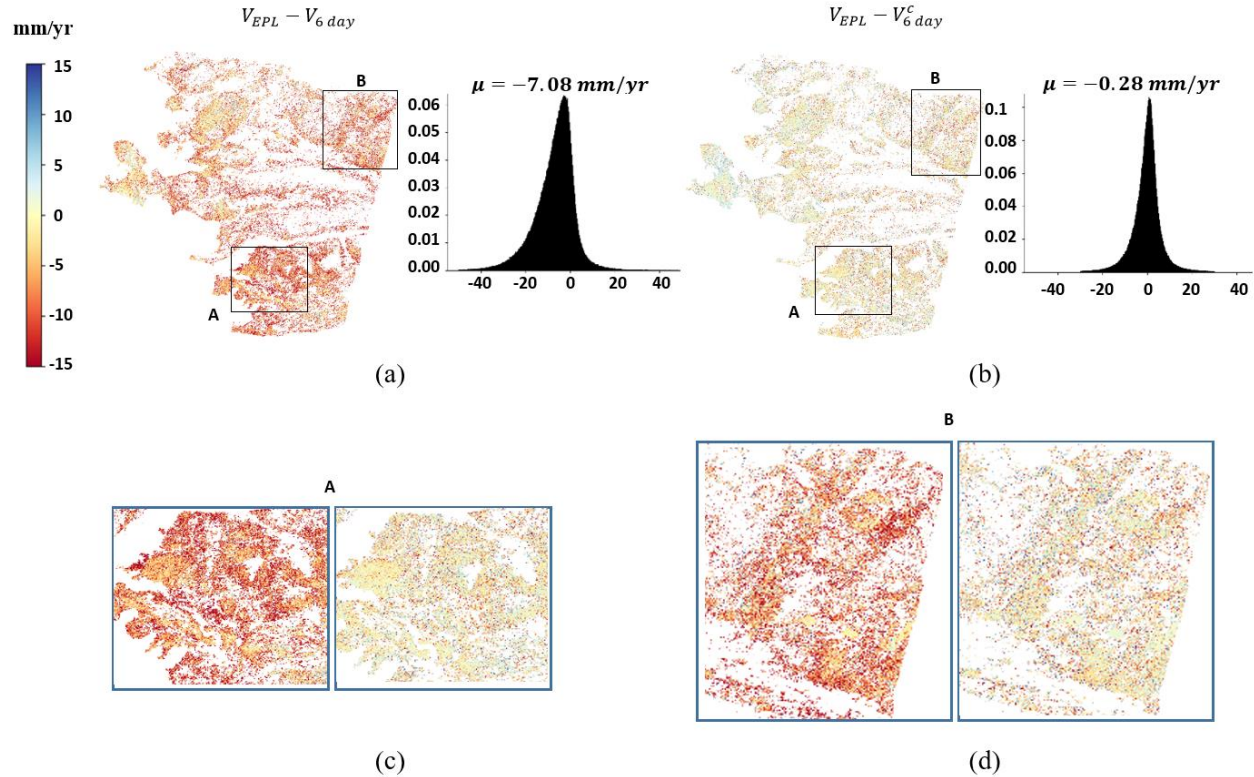
$$T = \sum_{i=1}^N \lambda_i v_i v_i^H \quad (14)$$

260 where the eigenvalues λ_i are arranged in descending order as $\lambda_1 \geq \lambda_2 \geq \dots \geq \lambda_N$, v_i is the
 261 corresponding eigenvector associated with eigenvalue λ_i and H stands for the conjugate transpose.

262 Phases $\hat{\phi}$ are estimated by extracting the phases of the eigenvector associated with the largest
 263 eigenvalue. The EPL velocity i.e. V_{EPL} can then be estimated using these linked phases. Full details
 264 of the algorithm are described in (De Zan et al. 2007; Fornaro et al. 2015). We used the a posteriori
 265 coherence of (Ferretti et al. 2011) as a quality measure for phase estimation. In this study, we chose
 266 a value of 0.4 as a threshold for the a posteriori coherence to mask out the unreliable phases.



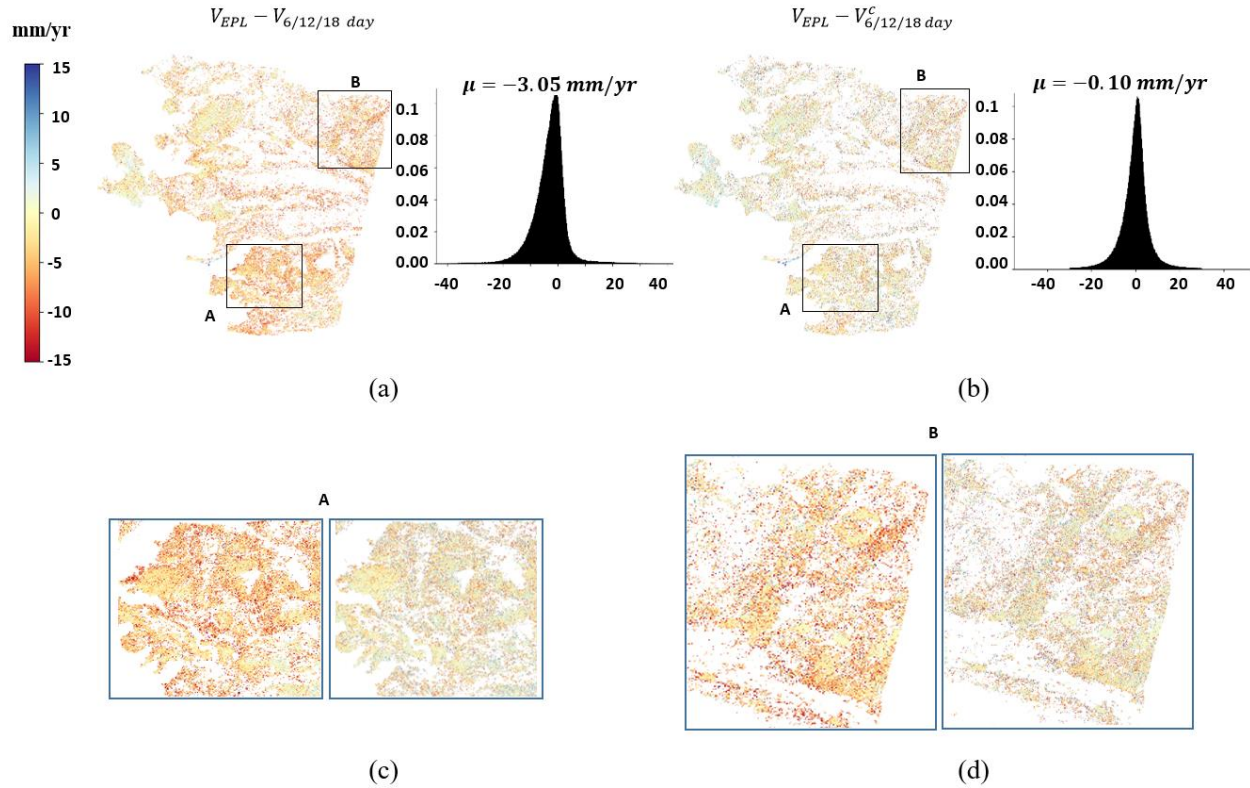
269 Figure 11) Scatterplot of 360-day velocities obtained from Eigen decomposition Phase Linking (V_{EPL}) against velocities obtained
 270 from (a) uncorrected 6-day interferograms, $V_{6\text{ day}}$, and (b) 6-day interferograms corrected with our empirical approach $V_{6\text{ day}}^c$



271 Figure 12) Effect of the phase bias correction on the velocity estimation. Difference between the EPL velocity and (a) the
 272 uncorrected 6-day velocity i.e. $V_{EPL} - V_{6\text{ day}}$, (b) corrected 6-day velocity i.e. $V_{EPL} - V_{6\text{ day}}^c$ are shown as maps and histograms.
 273 (c,d) zoomed-in views of the two subsets for areas A and B.

274 We calculated velocities from our corrected and uncorrected 6-day interferograms over the 360-
 275 day time period and checked the effectiveness of our bias correction strategy by comparing our
 276 estimated velocities before and after correction with the EPL velocities (Figures 11, 12).

277 A scatterplot of the original 6-day estimated velocities for all pixels in our Turkey frame, $V_{6\text{ day}}$,
 278 versus the velocities from EPL, V_{EPL} , is skewed to left (Figure 11 (a)), indicating that the velocities
 279 for many of the pixels in the uncorrected 6-day velocities have a negative bias. By comparison,
 280 the scatterplot of the corrected 6-day velocities, $V_{6\text{ day}}^c$, versus the V_{EPL} is centred on the diagonal
 281 1:1 line (Figure 11 (b)) indicating a high-degree of correlation between $V_{6\text{ day}}^c$ and V_{EPL} and a
 282 dramatic reduction in the phase bias. The coefficient of determination, R^2 , increases from 0.19



283

284 Figure 13. The effect of the phase bias correction on the velocity estimation. $V_{EPL} - V_{6/12/18 \text{ day}}$ is shown in (a) and $V_{EPL} -$
 285 $V_{6/12/18 \text{ day}}^c$ is shown in (b). The zoomed-in view of the two subsets A and B are shown in panels (c) and (d) respectively.

286 before the correction to 0.63 after correction, and the RMSE decreased from 11.3 to 6.28 after
 287 correcting for the interferograms.

288 Plotting maps of the difference between velocities obtained with EPL and those calculated from
 289 our corrected/uncorrected 6-day interferograms (Figure 12) confirms that the corrected velocities
 290 are much closer to those from EPL. Comparing the histograms shows that the mean and standard
 291 deviation of the differences between the velocities, changes from $-7.08 \pm 8.8 \text{ mm/yr}$, to
 292 $-0.28 \pm 6.2 \text{ mm/yr}$ after correction.

293 In the next experiment, we included all the 6, 12 and 18-day interferograms in our velocity
 294 estimation. We calculated velocities using both the original ($V_{6/12/18 \text{ day}}$) and the corrected
 295 interferograms ($V_{6/12/18 \text{ day}}^c$) and calculated their difference with V_{EPL} (Figure 13). Using the
 296 corrected interferograms decreases the mean velocity bias to -0.1 mm/yr .

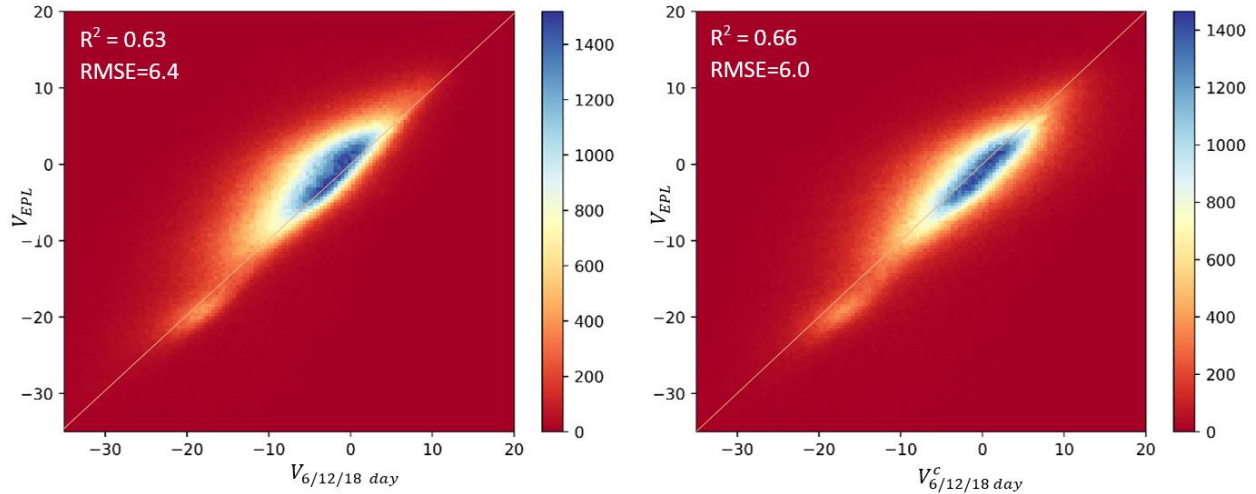
297 Similar to the velocity estimation using 6-day interferograms only, the scatterplot (Figure 14)
 298 shows good correlation between EPL velocities and those estimated from 6/12/18-day
 299 interferograms, after correction. The coefficient of determination, R^2 , increases from 0.63 before
 300 the correction to 0.66 after correction, and the RMSE decreased from 6.4 to 6.0 after correcting
 301 for the interferograms.

302 Table 1 shows a summary of the average velocities obtained with the 6-day and 6/12/18-day
 303 interferograms before and after correction in different land cover classes. The EPL estimated
 304 velocities are also given in this table. For all land cover classes, our corrected velocities agree well
 305 with those from phase linking approach.

306 Table 1) Summary of the average velocities in mm/yr obtained for all pixels and in different land covers

	ALL	URBAN	FOREST	CROPLAND
$V_{6 \text{ day}}$	10.8	7.1	15.0	40.2
$V_{6 \text{ day}}^c$	4.0	3.5	4.5	12.8
$V_{6/12/18 \text{ day}}$	6.7	5.0	8.8	21.1
$V_{6/12/18 \text{ day}}^c$	3.8	3.5	4.3	9.2
V_{EPL}	3.7	4.7	4.2	11.0

313



(a)

(b)

314 Figure 14) Scatterplot of 360-day velocities obtained from Eigendecomposition Phase Linking (V_{EPL}) against velocities obtained
 315 from (a) uncorrected 6/12/18-day interferograms, $V_{6/12/18 \text{ day}}$, and (b) 6/12/18-day interferograms corrected with our empirical
 316 approach $V_{6/12/18 \text{ day}}^c$
 317

318 Comparing the Figures 12 and 13 reveals that correcting the interferograms using the proposed
 319 strategy led to consistent phases. Clearly, upon correcting for the phase bias using the proposed
 320 method, it does not make any difference which stack of the interferograms be used for velocity
 321 estimation and the 6-day velocity will have very similar performance as the 6, 12 and 18-day
 322 velocity. This proves the consistency of the proposed strategy for correcting the phase bias.

323 7 Conclusions

324 We have shown that short-interval interferograms can be highly affected by phase bias (also known
 325 as fading signals) and the accumulation of this phase bias in time can highly affect the estimated
 326 velocity. We provide a readily applicable method to estimate the bias corrections for the
 327 interferograms. The proposed correction strategy is simple and effective in addressing the phase
 328 bias by providing a close performance to the phase linking approach. The method relies on the
 329 estimation of two constant regularization parameters, which can easily be calculated using a single
 330 long-term interferogram. The proposed method is based on the assumption that the phase bias in

331 an interferogram is linearly related to the sum of the bias shorter interferograms spanning the same
332 time. In this study, we used constant values for a_1 and a_2 , which relate the biases in 6-day
333 interferograms with those in 12-day and 18-day interferograms. Further investigation is needed to
334 determine if these are universal constants or if they vary spatially. We note that a similar approach
335 could also be developed for areas where the revisit time for Sentinel-1 is 12 days.

336 Though efficient and robust, PL approaches are computationally expensive both in terms of
337 generating $N(N - 1)/2$ interferograms and estimating the $(N - 1)$ optimal phases through often
338 iterative optimizations of the underlying covariance matrix. However, our proposed method only
339 requires calculating $(3N - 6)$ interferometric phases and solves for the bias correction using
340 through a single-step, straightforward and inexpensive least square inversion of Equation (10).
341 This is of high importance, particularly for automatic InSAR systems such as COMET-LiCSAR,
342 which are designed to automatically produce InSAR products by processing all Sentinel-1
343 acquisitions in a frame (~60 new 6-day acquisitions per frame per year over Europe).

344 More importantly, the quality of the PL estimated phases highly depends on the availability of the
345 long-term interferograms. In case of the decorrelated regions such as forest or agricultural areas,
346 where long-term coherence is difficult to maintain, the a posteriori coherence is degraded. Our
347 proposed method, on the other hand, is immune to this coherence loss as it only relies on the short
348 term interferograms (6/12/18 day in this study) for estimating the correction terms. We identified
349 a total of 2,400,000 points as coherent pixels (in section 5), whereas this number was decreased to
350 1,300,000 points when using the EPL approach (in section 6). Therefore, our correction method

351 can be applied to global compilations of short-term interferograms and offers the possibility of
352 accurate long-term velocities without a requirement for coherence in long-term interferograms.

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360 **References**

- 361
362 Ansari, H., Zan, F.D., & Bamler, R. (2018). Efficient Phase Estimation for Interferogram Stacks. *IEEE Transactions*
363 *on Geoscience and Remote Sensing*, 56, 4109-4125
- 364 Ansari, H., Zan, F.D., & Parizzi, A. (2021). Study of Systematic Bias in Measuring Surface Deformation With SAR
365 Interferometry. *IEEE Transactions on Geoscience and Remote Sensing*, 59, 1285-1301
- 366 Berardino, P., Fornaro, G., Lanari, R., & Sansosti, E. (2002). A new algorithm for surface deformation monitoring
367 based on small baseline differential SAR interferograms. *IEEE Transactions on Geoscience and Remote Sensing*,
368 40, 2375-2383
- 369 Biggs, J., Burgmann, R., Freymueller, J.T., Lu, Z., Parsons, B., Ryder, I., Schmalzle, G., & Wright, T. (2009). The
370 postseismic response to the 2002 M 7.9 Denali Fault earthquake: constraints from InSAR 2003–2005. *Geophysical*
371 *Journal International*, 176, 353-367
- 372 Cao, N., Lee, H., & Jung, H.C. (2016). A Phase-Decomposition-Based PSInSAR Processing Method. *IEEE*
373 *Transactions on Geoscience and Remote Sensing*, 54, 1074-1090
- 374 De Zan, F., & Gomba, G. (2018). Vegetation and soil moisture inversion from SAR closure phases: First
375 experiments and results. *Remote Sensing of Environment*, 217, 562-572
- 376 De Zan, F., Parizzi, A., Prats-Iraola, P., & López-Dekker, P. (2014). A SAR Interferometric Model for Soil
377 Moisture. *IEEE Transactions on Geoscience and Remote Sensing*, 52, 418-425
- 378 De Zan, F., Rocca, F., & Rucci, A. (2007). PS PROCESSING WITH DECORRELATING TARGETS. In
379 De Zan, F., Zonno, M., & López-Dekker, P. (2015). Phase Inconsistencies and Multiple Scattering in SAR
380 Interferometry. *IEEE Transactions on Geoscience and Remote Sensing*, 53, 6608-6616
- 381 Farr, T.G., Rosen, P.A., Caro, E., Crippen, R., Duren, R., Hensley, S., Kobrick, M., Paller, M., Rodriguez, E., Roth,
382 L., Seal, D., Shaffer, S., Shimada, J., Umland, J., Werner, M., Oskin, M., Burbank, D., & Alsdorf, D. (2007). The
383 Shuttle Radar Topography Mission. *Reviews of Geophysics*, 45
- 384 Ferretti, A., Fumagalli, A., Novali, F., Prati, C., Rocca, F., & Rucci, A. (2011). A New Algorithm for Processing
385 Interferometric Data-Stacks: SqueeSAR. *IEEE Transactions on Geoscience and Remote Sensing*, 49, 3460-3470
- 386 Fornaro, G., Verde, S., Reale, D., & Pauciuillo, A. (2015). CAESAR: An Approach Based on Covariance Matrix
387 Decomposition to Improve Multibaseline–Multitemporal Interferometric SAR Processing. *IEEE Transactions on*
388 *Geoscience and Remote Sensing*, 53, 2050-2065
- 389 Foroughnia, F., Nemati, S., Maghsoudi, Y., & Perissin, D. (2019). An iterative PS-InSAR method for the analysis of
390 large spatio-temporal baseline data stacks for land subsidence estimation. *International Journal of Applied Earth*
391 *Observation and Geoinformation*, 74, 248-258

392 Goldstein, R.M., & Werner, C.L. (1998). Radar interferogram filtering for geophysical applications. *Geophysical*
393 *Research Letters*, 25, 4035-4038

394 Guarnieri, A.M., & Tebaldini, S. (2008). On the Exploitation of Target Statistics for SAR Interferometry
395 Applications. *IEEE Transactions on Geoscience and Remote Sensing*, 46, 3436-3443

396 Juncu, D., Árnadóttir, T., Hooper, A., & Gunnarsson, G. (2017). Anthropogenic and natural ground deformation in
397 the Hengill geothermal area, Iceland. *Journal of Geophysical Research: Solid Earth*, 122, 692-709

398 Lazecký, M., Spaans, K., González, P.J., Maghsoudi, Y., Morishita, Y., Albino, F., Elliott, J., Greenall, N., Hatton,
399 E., Hooper, A., Juncu, D., McDougall, A., Walters, R.J., Watson, C.S., Weiss, J.R., & Wright, T.J. (2020). LiCSAR:
400 An Automatic InSAR Tool for Measuring and Monitoring Tectonic and Volcanic Activity. *Remote Sensing*, 12

401 Massonnet, D., Briole, P., & Arnaud, A. (1995). Deflation of Mount Etna monitored by spaceborne radar
402 interferometry. *Nature*, 375, 567-570

403 Michaelides, R.J., Zebker, H.A., & Zheng, Y. (2019). An Algorithm for Estimating and Correcting Decorrelation
404 Phase From InSAR Data Using Closure Phase Triplets. *IEEE Transactions on Geoscience and Remote Sensing*, 57,
405 10390-10397

406 Morishita, Y., Lazecky, M., Wright, T.J., Weiss, J.R., Elliott, J.R., & Hooper, A. (2020). LiCSBAS: An Open-
407 Source InSAR Time Series Analysis Package Integrated with the LiCSAR Automated Sentinel-1 InSAR Processor.
408 *Remote Sensing*, 12

409 Samiei-Esfahany, S., Martins, J.E., Leijen, F.v., & Hanssen, R.F. (2016). Phase Estimation for Distributed Scatterers
410 in InSAR Stacks Using Integer Least Squares Estimation. *IEEE Transactions on Geoscience and Remote Sensing*,
411 54, 5671-5687

412 Temtime, T., Biggs, J., Lewi, E., Hamling, I., Wright, T., & Ayele, A. (2018). Spatial and temporal patterns of
413 deformation at the Tendaho geothermal prospect, Ethiopia. *Journal of Volcanology and Geothermal Research*, 357,
414 56-67

415 Walters, R.J., Holley, R.J., Parsons, B., & Wright, T.J. (2011). Interseismic strain accumulation across the North
416 Anatolian Fault from Envisat InSAR measurements. *Geophysical Research Letters*, 38

417 Weiss, J.R., Walters, R.J., Morishita, Y., Wright, T.J., Lazecky, M., Wang, H., Hussain, E., Hooper, A.J., Elliott,
418 J.R., Rollins, C., Yu, C., González, P.J., Spaans, K., Li, Z., & Parsons, B. (2020). High-Resolution Surface
419 Velocities and Strain for Anatolia From Sentinel-1 InSAR and GNSS Data. *Geophysical Research Letters*, 47,
420 e2020GL087376

421 Zwieback, S., Liu, X., Antonova, S., Heim, B., Bartsch, A., Boike, J., & Hajnsek, I. (2016). A Statistical Test of
422 Phase Closure to Detect Influences on DInSAR Deformation Estimates Besides Displacements and Decorrelation
423 Noise: Two Case Studies in High-Latitude Regions. *IEEE Transactions on Geoscience and Remote Sensing*, 54,
424 5588-5601

425