

Tidal marsh resilience to sea level rise controlled by vertical accretion and landward migration under nature-based human adaptation scenarios

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Competing Interests

No competing interests need to be declared in relation to the work described.

Author Contributions

X.O. and S.Y.L. designed the study with R.M.C. contributing ideas to the study. X.O. collected, analysed the data and conducted mathematical modelling. Manuscript writing was led by X.O. with contributions from S.Y.L. and R.M.C.

1 **Abstract:** Tidal marshes are not only lost to human disturbance but also face the threat of sea
2 level rise (SLR). However, current earth system models used to estimate future changes in
3 wetland extent omit wetland's real responses to SLR without field observations. We
4 synthesised global data on sediment accretion rate (SAR) and surface elevation change (SEC)
5 for tidal marshes and developed a mathematical model to assess their resilience to future
6 SLR. Sediment loadings and precipitation largely explain the variance of marsh SAR and
7 SEC. Human disturbance resulted in less sediment accretion and existing conservation
8 activities were inefficient in promoting sediment accretion. Under the representative
9 concentration pathways and nature-based human adaptation scenarios, tidal marshes will gain
10 up to 63% of the current area by 2100 if sufficient sediment loadings and accommodation
11 space allow landward migration. If current accommodation space maintains, net areal losses
12 of > 30% are possible and hotspots of future marsh loss are largely in North America,
13 Australia and China. Projections for most SLR scenarios see marsh area peaking in the mid
14 rather than late 21st century. This implicates that tidal marshes may contribute to achieve a
15 climate neutral world by 2050. We highlight the importance of nature-based adaptation in
16 enhancing the resilience of tidal marshes to future SLR.

17 **Key Words:** Sea level rise; Nature-based human adaptation; Human disturbance; Climate
18 neutral; Sediment accretion; Surface elevation change

19 **1 Introduction**

20 Tidal marshes are widely distributed 'blue carbon' ecosystems, providing substantial
21 ecosystem services (e.g. coastal protection, water purification and nursery ground for
22 fisheries) (Barbier 2012) valued at US\$194,000 ha⁻¹ yr⁻¹ (Costanza et al., 2014). Nonetheless,
23 the area of marshes has been declining globally due to direct adverse anthropogenic impacts
24 (Craft et al., 2008; Doody 2004; Gedan et al., 2009; Koch et al., 2009) and wind-wave
25 erosion of marsh edges (Leonardi et al., 2018; Marani et al., 2011), the effect of which on

26 marsh loss has not been accounted for. Tidal marshes now also face an even more perilous
27 future because of sea level rise (SLR). Loss of tidal marshes to SLR is partly mitigated by
28 vertical sediment accretion, resulting from the balance of below-ground root production and
29 decomposition (Ouyang et al., 2017), and sediment input from marine (Chmura and Hung,
30 2004) and/or riverine sources (Craft 2007).

31

32 Further, there are feedbacks and interactions among plant growth, geomorphology and
33 hydrodynamics that allow tidal marshes to resist submergence due to SLR (Gedan et al.,
34 2011; Kirwan and Megonigal, 2013; Marani et al., 2006). Tidal marsh sediments also show
35 different deposition patterns dependent on morphodynamics, which shape tidal marshes
36 through combined ecogeomorphology and hydrodynamics (Friedrichs and Perry, 2001). In
37 addition to sediment accretion, tidal marsh surface elevation change depends on erosion and
38 subsurface processes, including compaction, shrink-swell, subsidence, decomposition and
39 tectonic adjustments (Cahoon et al., 2011; Rogers et al., 2006). Sediment accretion may
40 alleviate the impact of SLR on tidal marshes when surface elevation increases at or exceeds
41 the rate of local relative SLR (Aniesfeld et al., 2016). Conversely, tidal marshes could be
42 submerged (Andersen et al., 2011) and drowned, compromising C sequestration and other
43 services (Ouyang and Lee, 2014). Potentially, massive losses of tidal marshes are possible, in
44 line with that described for mangroves across the Indo-Pacific region, i.e. more than half the
45 mangrove forests have already lost elevation relative to sea level (Lovelock et al., 2015).

46

47 Sediment accretion and surface elevation in tidal marshes are regulated by biotic and abiotic
48 factors (including climatic, geographic, tidal, and local factors such as supply of suspended
49 sediment). Sediment accretion may also vary over different time scales (Breithaupt et al.
50 2018). While these factors likely drive the change in sediment accretion and/or surface

51 elevation change, their relative influences on the global pattern of tidal marsh sediment
52 accretion is poorly known. An understanding of this pattern can substantially improve
53 projections of tidal marsh response to SLR in future earth system models.

54

55 A recent study of the vulnerability of regional mangroves to SLR revealed the importance of
56 time scales in evaluating coastal wetland vertical change in response to SLR (Breithaupt et al.
57 2018). The two distinct ecotypes of coastal wetlands, i.e. mangroves and tidal marshes,
58 however, are different in nature, e.g. distribution (tropical and subtropical vs. global) and
59 plant traits (mainly woody vs. shrubby/ herbaceous). Further, recent evaluations of the
60 vulnerability of coastal wetlands to SLR are mainly derived from SEC and short-term
61 sediment accretion rate (SAR) (Lovelock et al., 2015; Kirwan et al., 2016) or calibrate their
62 results using short-term SAR (Schuerch et al., 2018), but lack long-term rates estimated by
63 methods such as radiometric geochronology (Parkinson et al., 2017). The short-term records
64 are very unlikely to incorporate the cumulative effects of subsurface processes (Parkinson et
65 al., 2017), e.g. root growth and compaction (Allen 2000; Rybczyk and Cahoon, 2002), which
66 also strongly influence sediment accretion (Kolker et al., 2009). We aim to synthesise the
67 global data to properly assess the important relationship between sediment accretion/
68 elevation change and their potential drivers (See Table 1 for comparison of studies on the
69 response of coastal wetlands to SLR).

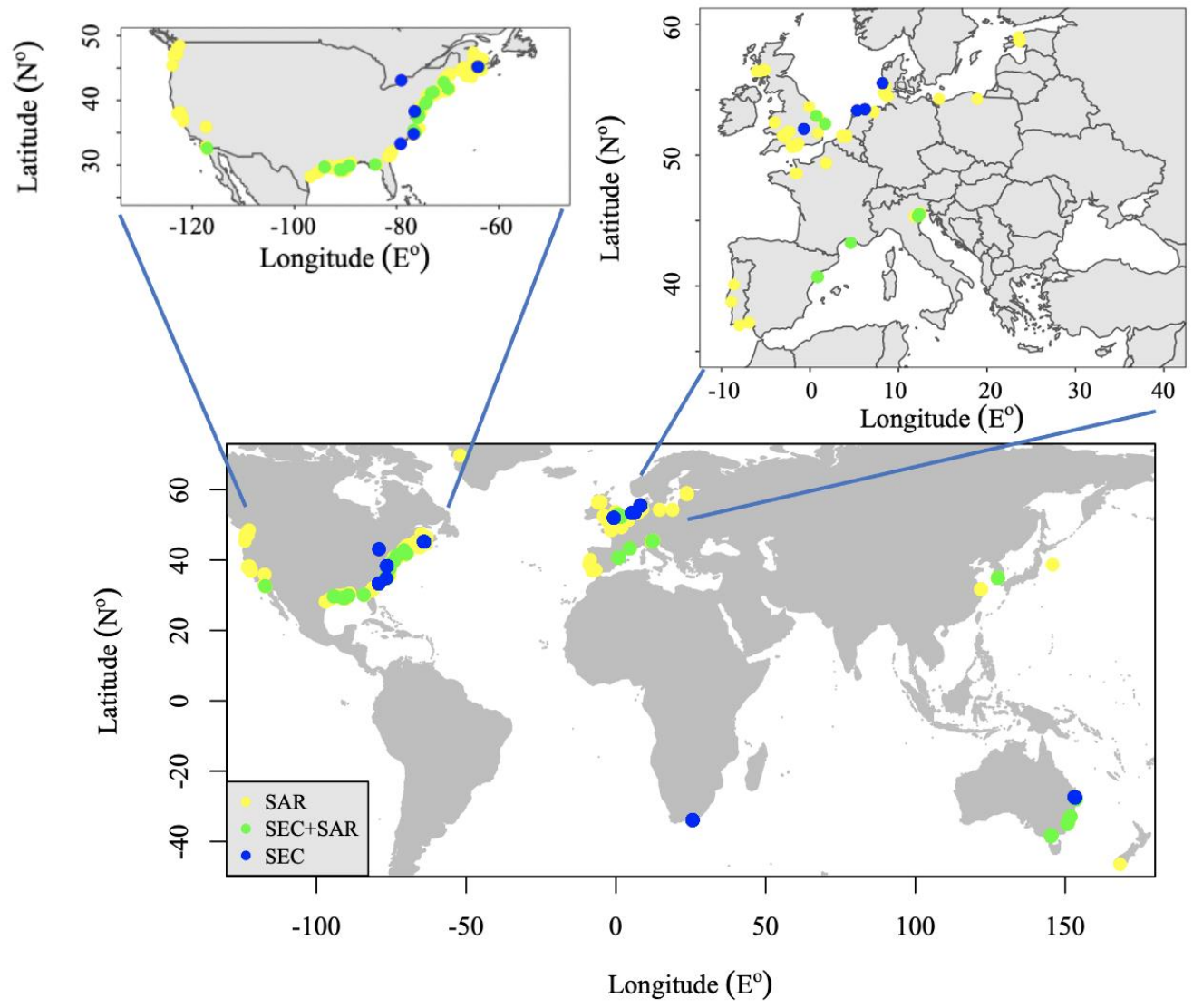
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71 We reviewed data on both short- and long-term global tidal marsh SAR and SEC (Fig. 1a),
72 and provide the first systematic test of the quantitative influence of potential drivers on global
73 tidal marsh SAR and SEC. We demonstrate that both sediment supply (Törnqvist et al., 2019)
74 and landward migration contributing to elevation capital (Schuerch et al., 2018, 2019) are
75 important, and predict the resilience of tidal marshes to future SLR by establishing an

76 Table 1 A comparison of studies evaluating the response of coastal wetlands to sea level rise

| Meta-analysis/Quantitative reviews | | | | |
|---|--|---|--------------------|------------------------------------|
| Land building via vertical sediment accretion | Lateral landward migration via accommodation space | Controls on sediment accretion/surface elevation change | Geographic regions | References |
| Using real data on surface elevation change to assess the response of mangroves to sea level rise | Not considered | Assessed the impact of controls such as climatic, geomorphic, geographic factors on mangrove sediment accretion/surface elevation change | Indo-Pacific | Lovelock et al. ²¹ |
| Comparing real data on sediment accretion and relative sea level rise to show marsh submerging or aggrading without considering shallow or deep subsidence | Not considered but stated the limitation | Not considered | Global | Kirwan et al. ²³ |
| Approximate land building using a wetland adaptation score that relies on suspended sediment concentration to describe sediment supply, which is conflated with the former | Assessed accommodation space available for wetland landward migration | Not considered | Global | Schuerch et al. ²⁴ |
| Using real data on surface elevation change to model the response of tidal marshes to sea level rise | Assessed accommodation space available for tidal marsh landward migration under human adaptations | Assessed the impact of controls such as human disturbance, climatic, geomorphic, geographic factors on marsh sediment accretion/surface elevation change | Global | This study |
| Qualitative reviews | | | | |
| Contributions | | Geographic regions | | References |
| Described historical sea level fluctuations in the region, sediment supply to maintain marsh elevation change, and controls on accommodation space for marsh migration | | Atlantic and southern North Sea coasts of Europe | | Allen ²⁶ |
| Stated whether wetlands continue to survive sea level rise depends largely on how human impacts interact with rapid sea level rise, and socio-economic factors that influence transgression into adjacent uplands | | NA | | Kirwan and Megonigal ¹³ |

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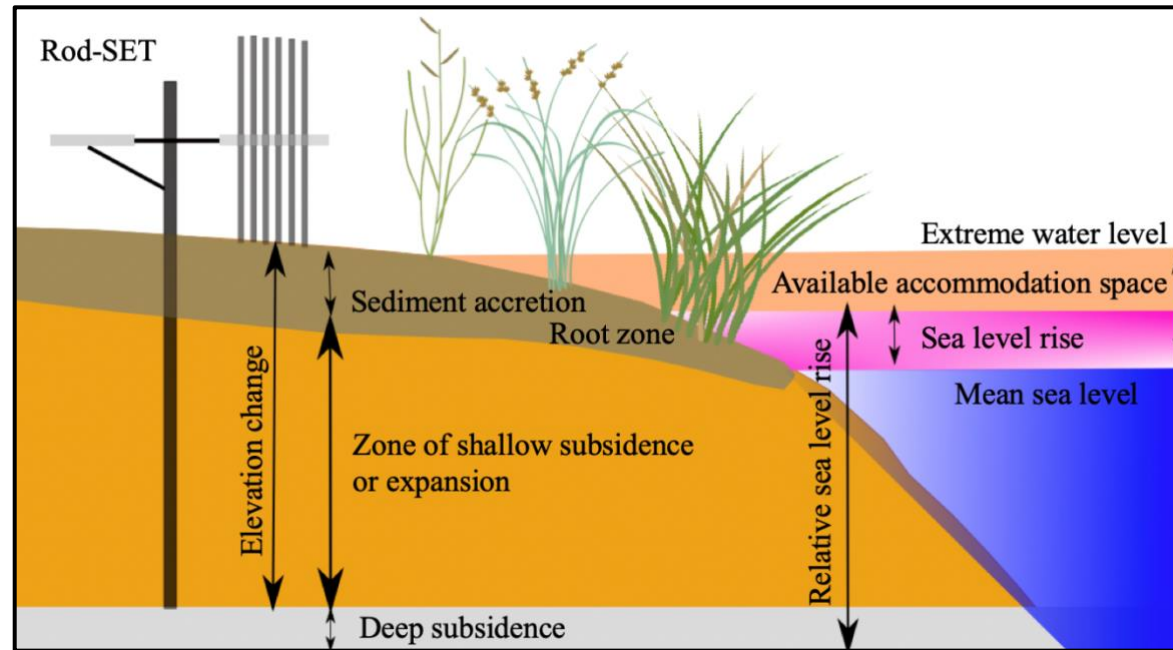
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103 **Fig. 1** Global distribution of study sites from the collated studies and a conceptual model showing marsh response to sea level rise. (a) sites of
104 SAR and/or SEC; (b) tidal marshes responding to SLR via vertical sediment accretion and accommodation space available for lateral migration.
105 The upper sub-panels in a) show the sites in North America and Europe, respectively. SAR and SEC denote sediment accretion rate and surface
106 elevation change, respectively. SET denotes surface elevation table.

107

108

109 integrated mathematical model incorporating the effects of vertical sediment accretion
110 described by SEC and lateral landward migration.

111

112 **2 Methods**

113 **2.1 Literature review**

114 We collected references on SAR and SEC in tidal marshes from
115 <http://www.sciencedirect.com/> and <http://pcs.webofknowledge.com/>. We collated 156
116 references containing 1229 independent measurements of SAR and 146 measurements of
117 SEC (Fig. 1a). SAR is the vertical dimension of sediment development determined from
118 short- and/or long-term markers, and integrates the sedimentological and biologic processes
119 contributing to organic and inorganic matter deposited on the sediment surface (Cahoon et
120 al., 1995; Thom 1992). SEC is the change in elevation relative to a subsurface datum, the
121 depth of which is determined by the technique used (Thom 1992). SAR results from sediment
122 accretion from allochthonous and autochthonous sources, while SEC results from not only
123 sediment accretion but also other processes (Fig. 1b) such as subsidence and autocompaction
124 (Rogers et al., 2006). The data cover tidal marshes over a latitudinal range from 69.7°N to
125 46.5°S. Besides SAR and SEC data, tidal range and tidal frequency at each sampling site
126 were collected from data reported in relevant references. Where data are unavailable, tidal
127 ranges were collected from the data at the nearest tidal stations, reported by National Oceanic
128 and Atmospheric Administration, USA (<http://tidesandcurrents.noaa.gov>). Tidal frequencies
129 were extracted from the map of world tidal patterns (Gabler et al., 2007), which are divided
130 into diurnal, semi-diurnal and mixed tides. Tidal ranges were sorted into micro- (<2 m),
131 meso- (2-4 m) and macro- (>4 m) tides. Marsh locations, geomorphology, sampling periods
132 and methods were also extracted from the references. Marsh locations were reported as levee
133 or back-marsh in our collected references. The former is close to the river/creek bank while

134 the latter is located behind the former at a higher tidal elevation. Geomorphology was sorted
135 into back-barrier, fluvial, transitional and bluff-toe marshes (Kelley et al., 1988). Species
136 richness were categorised as 1 or >1, since some studies report the occurrence of mixed
137 species and generally there are few studies at sites with three or more species. Plant types
138 were divided into herb or herb + shrub, where herbaceous plants co-occur with shrubs. While
139 SAR for shrubs were also reported, the number of studies is too few to allow a meaningful
140 comparison, as are SEC data for all plant types. Fourteen methods were used to estimate SAR
141 in our collated references, with time scales ranging from sub-decadal (e.g. several years),
142 decadal, to millennial, mainly assessed by radiometric isotopic dating (e.g. ²¹⁰Pb, ¹³⁷Cs and
143 ¹⁴C) and marker horizons (Supplementary Material Fig. S1). SAR at millennial scales was
144 estimated from ¹⁴C. SEC was estimated by the surface elevation table method.

145

146 **2.2 Satellite data collection**

147 We also collected data on precipitation, sea ice and total suspended matter, which are
148 possible drivers of SAR and SEC. Monthly precipitation data were extracted from the
149 gridded monthly total precipitation of Global Precipitation Climatology Centre, with a spatial
150 resolution of 0.5×0.5 degree latitude by longitude and time periods between 1901 and 2013,
151 and 2014 to date. Where data are unavailable, precipitation data were supplemented and
152 extracted from the gridded monthly precipitation of British Atmospheric Data Centre. Sea ice
153 data were extracted from the monthly gridded sea ice data set of the Met Office Hadley
154 Centre, with a spatial resolution of 1×1 degree latitude by longitude and time period between
155 1870 and 2017. The sea ice data are satellite microwave-based, compensating for the impact
156 of surface melt effects on retrievals in the Arctic and for algorithm deficiencies in the
157 Antarctic, and considering consistency with the historical *in situ* concentrations (Rayner et
158 al., 2003). Level-3-processed suspended matter data binned monthly (available from

159 <http://hermes.acri.fr/>) were used in our study. Suspended matter was extracted from the
160 GlobColour primary data set of the European Space Agency's Envisat satellite (390 - 1,040
161 nm), which is built using the Medium Resolution Imaging Spectrometer and provides total
162 suspended matter of marine waters between 2002 and 2011 with a spatial resolution of four
163 km. Data products were processed and validated as part of the European Space Agency's Data
164 User Element GlobColour Global Ocean Colour for Carbon Cycle Research project (refer to
165 http://www.globcolour.info/CDR_Docs/GlobCOLOUR_PUG.pdf for more information about
166 data processing). Suspended matter in coastal waters is an indicator of river-runoff and
167 resuspension as well as sediment loadings to coastal wetlands. This indicator is useful in both
168 estuarine and reef lagoon waters (Blondeau-Patissier et al., 2014) and has been used as an
169 estimate for mangrove SEC (Lovelock et al., 2015).

170

171 **2.3 Analysis of controls on SAR and SEC**

172 We examined the relationship between SAR/SEC and influential factors, including latitude,
173 precipitation, sea ice concentration, total suspended matter, tidal range and tidal frequencies
174 in a statistical model. The categorical variables are number coded, with one level of each
175 variable selected as the reference. The influences of other factors (e.g. species richness, plant
176 type, location, and elevation) on SAR and/or SEC were not included in the model due to
177 limited available data, which would have substantially reduced the overall degrees of
178 freedom if included. We started the model with multiple regression, but the hypothesis of
179 normality was grossly violated and the residuals were highly heterogeneous. The relative
180 influence of explanatory variables on SAR and SECs was analysed via boosted regression
181 tree (BRT) models. The models were established with a tree complexity of 5, learning rate of
182 0.005, back fraction of 0.5 and 10-fold cross-validation optimization. Gaussian distribution
183 was set for SAR and SECs (Supplementary Material Fig. S2). We also examined the

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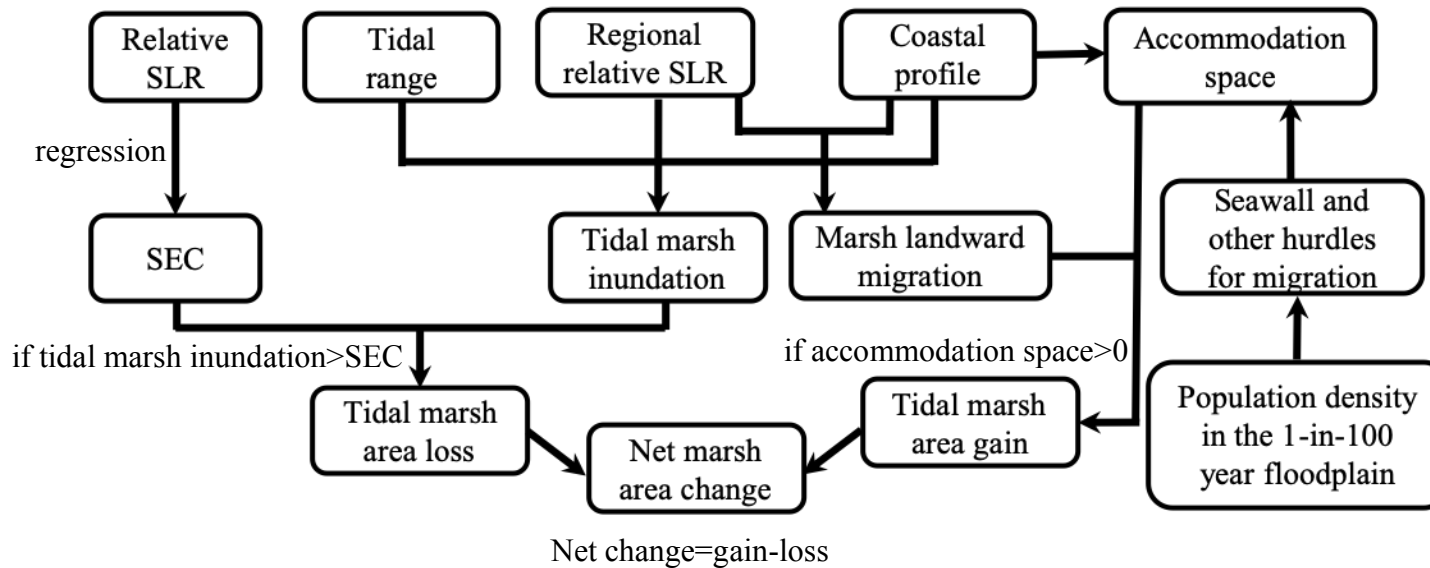
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192 **Fig. 2** A diagram describing the integrated mathematical modelling framework for simulating the response of tidal marshes to sea level rise. SLR
 193 denotes sea level rise. SEC denotes surface elevation change.

194 relationship between SEC and SAR, and the relationship between SEC/SAR and total
195 suspended matter, using a linear regression.

196

197 In addition to the above analyses, we examined the difference in SEC/SAR among
198 categorical variables that were significant in the BRT models, and other factors using the
199 Kruskal-Wallis rank sum test. Where significant differences were found, non-parametric
200 Mann-Whitney U tests were used to identify significant differences among the groups. Paired
201 t tests were used to compare the difference in SAR from the same sites at different time
202 scales (i.e. sub-decadal, decadal, centennial and millennial), or from the same sites at short,
203 mid- and long terms when the temporal span is different at the same time scales. These
204 univariate analyses address how SAR/SEC are driven by the categorical variables that could
205 not be addressed in the BRT models.

206

207 **2.4 Modelling the response of tidal marshes to SLR**

208 We evaluated the response of tidal marshes to future SLR using an integrated mathematic
209 model, which considers both the area gain and loss for tidal marshes (See the mathematic
210 modelling framework described in Fig. 2). Marsh area gains through the accommodation
211 space available for marsh landward migration, while area losses occur due to inundation if
212 marsh SEC cannot keep pace with RSLR, which is the resultant of SLR and other processes
213 such as subsidence and tectonic activity.

214

215 Landward migration of tidal marshes is facilitated by the available accommodation space
216 between mean seal level and mean high water spring. Marsh landward migration is likely
217 obstructed by sea defences and other man-made infrastructures such as seawalls, aquaculture
218 ponds and roads (Leonardi et al., 2018; Ma et al., 2014). Nonetheless, the global data on

219 infrastructure inhibiting marsh landward migration is lacking. We adopted the method of
220 Schuerch et al. (2018) to approximate accommodation space via different population density
221 thresholds, above which no accommodation space is considered available for tidal marshes to
222 migrate landward. The population density thresholds include population density of 5, 20, 150
223 and 300 people km⁻², with lower and upper population boundaries corresponding to nearly
224 uninhabited land (Mittermeier et al., 2003) and the European Commission's definition of
225 urban areas (Dijkstra and Poelman, 2014), respectively. A 5-year-lag is assumed for wetland
226 establishment when tidal marshes migrate landward, as evidenced by restoration practice
227 (Wolters et al., 2008).

228

229 With increasing sea levels, both the upper and lower boundaries of marsh migration shift,
230 likely resulting in marsh inundation if marsh SEC cannot keep pace with RSLR. We
231 estimated global SEC in tidal marshes from RSLR (precision: 20km) since our dataset cannot
232 cover all the tidal marshes. Specifically, we used the marsh SEC data collated from the
233 references and associated RSLR data under different SLR scenarios to establish the
234 relationships between marsh SEC and RSLR for different scenarios. The established
235 relationships were used to estimate SEC of global tidal marshes, which was combined with
236 RSLR and the coastal profile to determine marsh loss to inundation. The SLR scenarios
237 include the stringent mitigation scenario (RCP 2.6), intermediate scenarios (RCP 4.5), and
238 the scenario with very high greenhouse gas emissions (RCP 8.5) (IPCC 2014). Coastal
239 profiles were created by dividing the floodplain areas (Hinkel et al., 2014) per elevation
240 increment (from <1.5 m to 16.5 m at eight intervals) by the length of the corresponding
241 coastal segment to calculate the inundation lengths. Linear interpolation between the mean
242 high-water spring and an elevation of 1.5 m (or higher) in the coastal profile was used to
243 approximate high resolution light detection and range (LiDAR) derived elevations, with an

244 error of less than 30 cm (Titus and Wang, 2008). Net marsh area change is the difference
245 between marsh area gain and loss.

246

247 Statistical analyses were conducted using the R programming language (R Core Team, 2014).
248 Precipitation, sea ice and total suspended matter data were extracted by the package 'ncdf4'
249 (Pierce 2017). The packages 'dismo' (Hijmans et al., 2017) and 'gbm' (Greenwell et al., 2019)
250 were used to undertake BRT. The R packages 'rgeos' (Bivand and Rundel, 2020) and others
251 were used in the spatial analysis. The integrated mathematical modelling was performed in
252 Matlab R2015b.

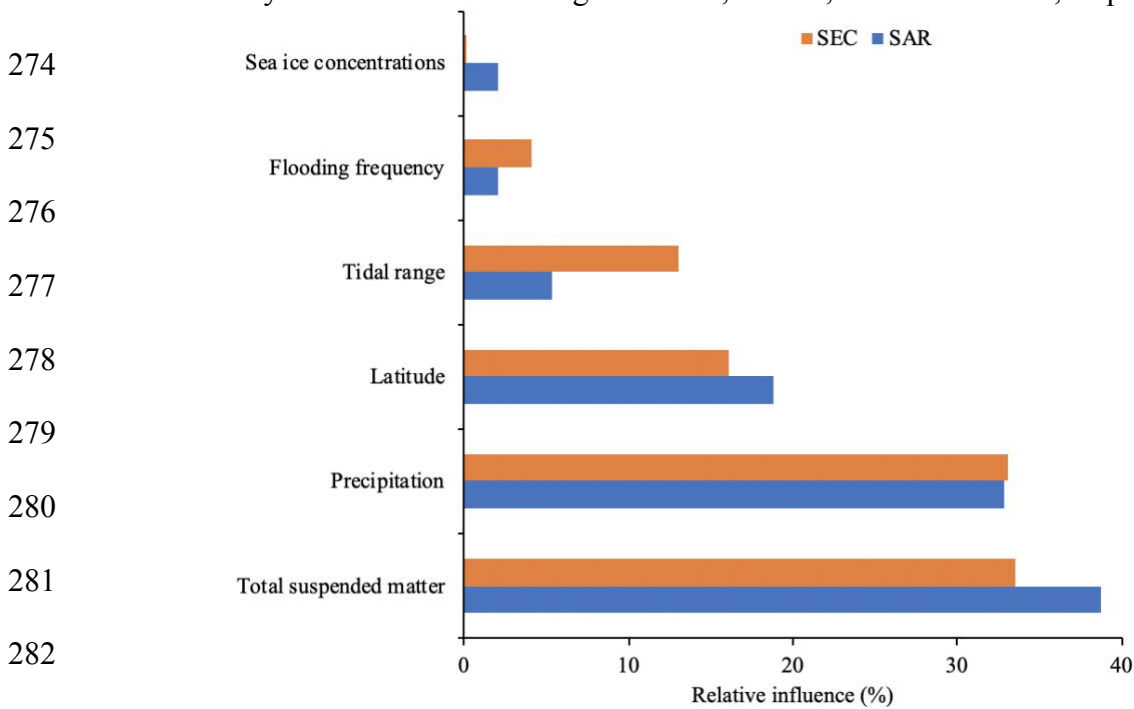
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254 **3 Results**

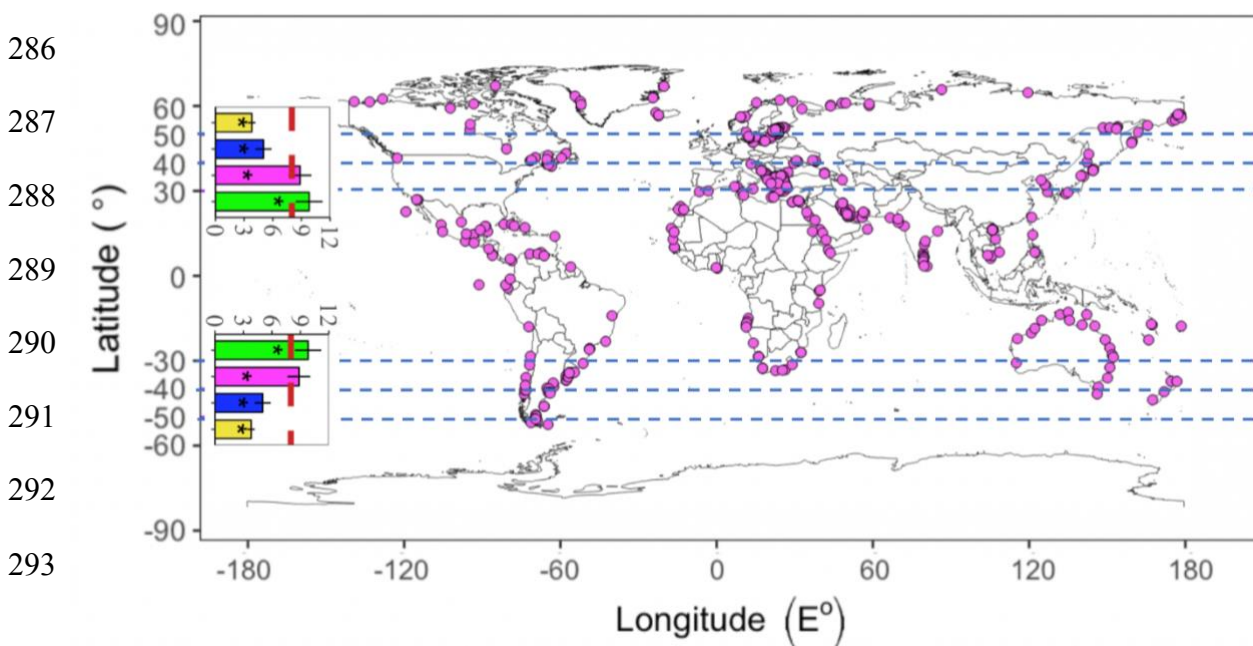
255 **3.1 Controls on tidal marsh SAR and SEC**

256 Boosted regression tree (BRT) modelling was used to explore the relative influence of
257 potential drivers on SAR and SEC (Supplementary Material Table S1). The cross-validation
258 procedure shows that the percentage of variance explained by the models are (mean, SE)
259 30.5% (2.7%) and 51.3% (14.1%) for SAR and SEC, respectively. The best-fit model shows
260 that SAR is driven by suspended matter, precipitation, latitude, tidal range, tidal frequency
261 and sea ice (Fig. 3). The model suggests that suspended matter and precipitation are the
262 dominant drivers and together account for 71.6% of the relative influences on SAR. The best-
263 fit model for SEC shows that it is driven by precipitation, suspended matter, latitude, tidal
264 range and tidal frequency. Again, precipitation and suspended matter are the dominant
265 drivers and account for 66.6% of the relative influences on SEC. Further exploration shows
266 that there are significant relationships between both SAR/SEC and suspended matter
267 (Supplementary Material Fig. S3a and b). There is a strong positive relationship between SEC
268 and SAR (Supplementary Material Fig. S3c, $R^2=0.82$, $p<<0.001$, F test). Based on this

269 relationship, we estimated surface elevation change in global saltmarshes (Fig. 4). The
 270 median values of surface elevation changes were used as the representative values of each
 271 latitudinal interval since both raw and transformed data did not meet the normal distribution
 272 assumption. Globally, the representative saltmarsh surface elevation changes are 5.9, 2.7, 2.3
 273 and 2.2 mm yr⁻¹ at the latitudinal ranges of <30°, 30-40°, 40-50° and >50°, respectively.



283 **Fig. 3** Relative influences of the explanatory variables on saltmarsh sediment accretion rates
 284 and surface elevation change. The orange and blue bars represent the relative influences on
 285 sediment accretion rates (SAR) and surface elevation change (SEC), respectively.

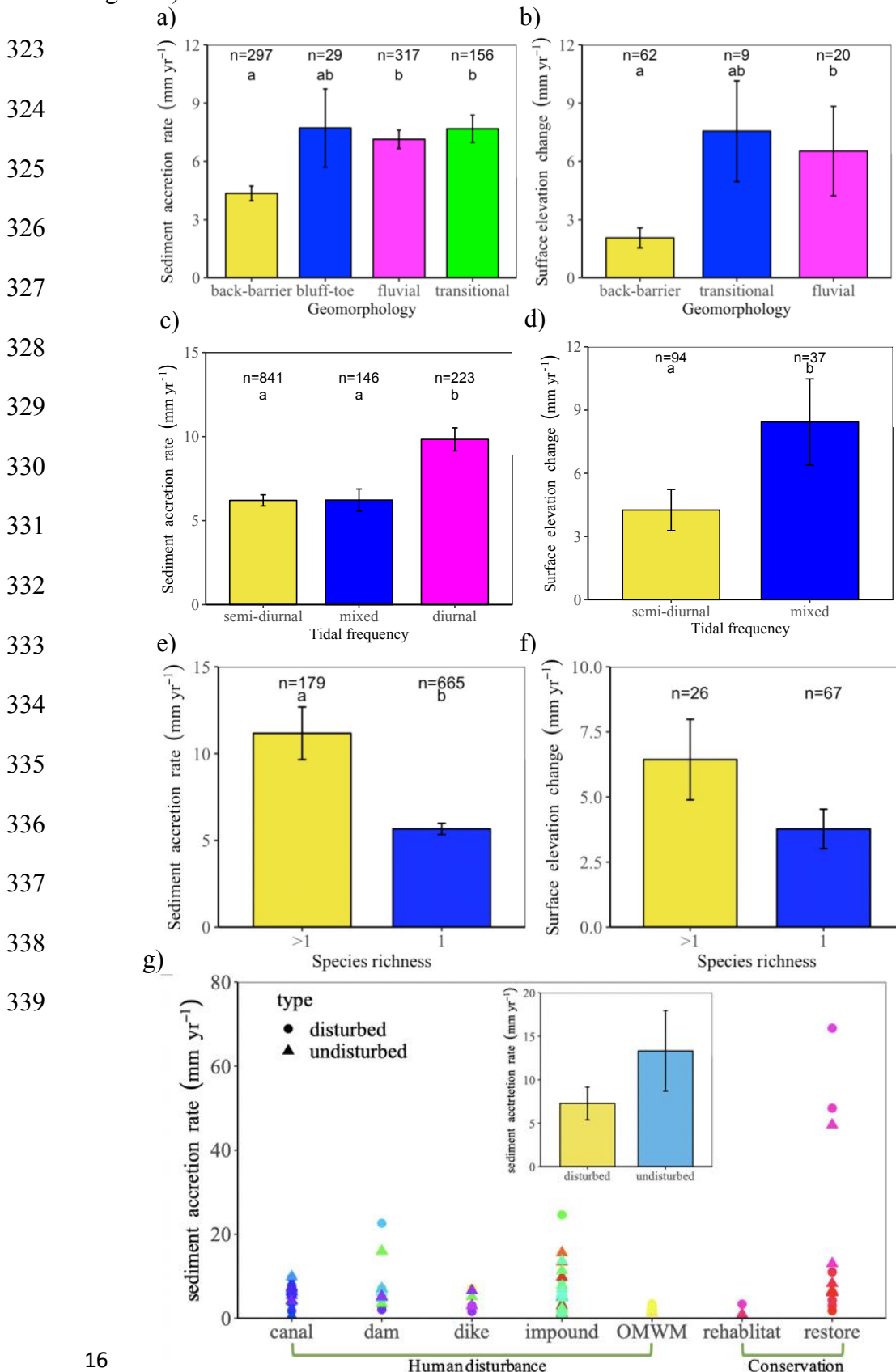


294 **Fig. 4** Surface elevation change of global saltmarshes at latitudinal ranges. The background
295 figure shows the global distribution of saltmarshes, depicted by pink dots. The error bar
296 charts show the estimated SEC (mm yr⁻¹) based on observations of sediment accretion rates at
297 latitudinal ranges <30°, 30-40°, 40-50° and >50° (n=198, 275, 451 and 283, respectively), as
298 denoted by the blue dashed lines. The stars in each error bar denote the representative SEC at
299 each latitudinal interval. The red dashed lines in the error bar charts are the lower limit of
300 SLR for RCP 8.5 at medium confidence to 2100.

301

302 We examined how SAR and SEC are driven by tidal frequencies and other potential drivers
303 that were excluded in the above models owing to limited data but enough for univariate
304 analyses. There are significant differences in both SAR and SEC among tidal marshes
305 inundated at different frequencies (Fig. 5c and d), and among geomorphologic settings (Fig.
306 5a and b). In particular, SAR in tidal marshes flooded by diurnal cycles are significantly
307 higher than those flooded by both semi-diurnal (Wilcoxon rank sum test, W=46300,
308 p<<0.001) and mixed cycles (W=9372, p<<0.001), while the latter two are not significantly
309 different. However, SECs in tidal marshes flooded by mixed tides are significantly higher
310 than those with semi-diurnal cycles (W=1235, p=0.01), but both show significant variances
311 due to the limited data. SAR are significantly higher in fluvial and transitional than back-
312 barrier marshes (W=32904, p<<0.001; W=16924, p<<0.001), while those of bluff-toe
313 marshes are not different from all others. SECs are significantly higher in fluvial than back-
314 barrier marshes (W=386, p=0.012), while that of transitional marshes are not different from
315 other marsh types. SAR varies significantly with species richness (W = 47690, p<<0.001,
316 Fig. 5e) and plant types (W = 20133, p=0.0028, Supplementary Material Fig. S4a), while
317 SECs do not vary significantly with species richness (W = 778, p=0.43, Fig. 5f). The
318 difference may lie in the limited observations on SEC, which show large variances. SAR at

319 sites with two or more species ($11.20 \pm 1.52 \text{ mm yr}^{-1}$, mean \pm SE) are almost double those at
 320 monospecific sites ($5.62 \pm 0.32 \text{ mm yr}^{-1}$). In addition, there are significant differences in
 321 sediment accretion rates among locations ($W = 7179$, $p=0.0004$, Supplementary Information
 322 Fig. S4b).



340 **Fig. 5** Variation of sediment accretion rate and/or surface elevation change with natural and
341 anthropogenic factors. (a) and (b), geomorphology; (c) and (d), tidal frequency; (e) and (f),
342 species richness; (g) human activities. Human activities include human disturbance and
343 conservation activities. OMWM denotes open marsh water management. Different sites in (g)
344 are depicted by different colours. Significance values are given in the text.

345

346 **3.2 Temporal changes in sediment accretion rates**

347 Sediment accretion rates fluctuate over different time scales, ranging from sub-decadal to
348 centennial scales. There is significant difference in sediment accretion rates between decadal
349 and sub-decadal scales (Supplementary Information Fig. S5, Paired t test, $t=2.2$, $p=0.033$,
350 $df=36$). For sediment accretion rates at the same temporal scale but different temporal span,
351 differences between mid- and short centennial scales are significant ($t=-4.2$, $p=0.002$, $df=9$).
352 This result further corroborates our conclusion that assessments of the vulnerability of coastal
353 wetlands to sea level rise with only short-term sediment records are inappropriate.

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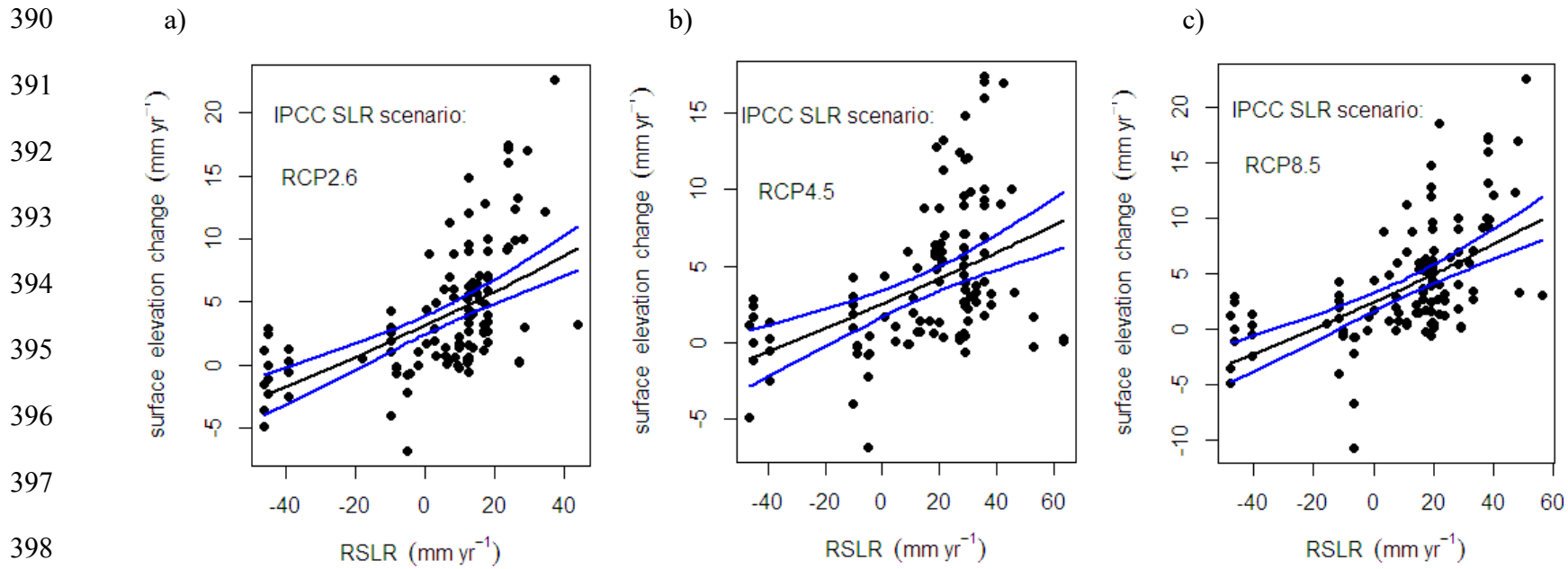
355 **3.3 The impact of human activities on SAR**

356 The impact of different human activities on SAR was compared for sites where conservation
357 and/or human disturbances are reported in conjunction with sediment accretion. SAR are
358 significantly lower in human-disturbed than in undisturbed sites (Fig. 5g, paired t test, $t=-$
359 2.02 , $p=0.048$, $df=50$), while in-situ conservation (e.g. inclusion of tidal marshes in reserves)
360 had no significant impact on SAR, probably because of the significance of ex-situ drivers,
361 e.g. allochthonous sediment supply. Among sites affected by human disturbances,
362 impoundment is the most frequently reported activity. SAR at impounded sites are
363 significantly higher than those of natural sites (Fig. 5g, paired t test, $t=-2.1$, $p=0.047$, $df=23$).

364

365 **3.4 The resilience of tidal marshes to SLR**

366 Significant relationships exist between tidal marsh SEC and RSLR under SLR scenarios of
367 SLR RCP2.6 ($R^2=0.33$, $p<0.05$, F test), RCP4.5 ($R^2=0.21$, $p<0.05$, F test) and RCP8.5
368 ($R^2=0.34$, $p<0.05$, F test) (Fig. 6). These relationships are significant to estimating global
369 SEC since we cannot collect SEC from all the tidal marshes while global RSLR data are
370 available at relatively high precisions. The relationships were incorporated into the modelling
371 framework (Fig. 2) to simulate the response of tidal marshes to SLR. We estimated global
372 tidal marsh areal change by 2100 at decadal intervals under the representative concentration
373 pathways (RCPs) of global SLR (RCP2.6: 29 cm, RCP4.5: 50 cm and RCP8.5: 110 cm) and
374 thresholds of population density (5, 20, 150 and 300 people km^{-2} , Fig. 7). The thresholds of
375 population density correspond to the lower and upper boundaries of three human adaptation
376 scenarios: (1) a high level of nature-based adaptation with population density thresholds of
377 150-300 people km^{-2} in the 1-in-100-year coastal floodplain; (2) a moderate level of nature-
378 based adaptation with population density thresholds of 20-150 people km^{-2} ; and (3) a
379 business-as-usual scenario with population density thresholds of 5-20 people km^{-2} . Nature-
380 based adaptation at high levels corresponds to urban areas where more accommodation space
381 can be created while that at low levels corresponds to nearly uninhabited land, where
382 solutions are less necessary to create accommodation space. Human adaptation scenarios 1 to
383 3 correspond to a decrease in additional accommodation space created by human through
384 nature-based adaptation solutions. Globally, under human adaptation scenario 1, tidal marsh
385 areal change fluctuates between -7% and -5.3% for RCP2.6, and between 58.2% and 62.7%
386 for RCP8.5. Under scenario 2, tidal marsh areal change fluctuates between -24.4% and -7%
387 for RCP2.6, and between 36.3% and 58.2% for RCP8.5. Scenario 3 would see tidal marsh
388 areal change fluctuate between -30.8% and -24.4% for RCP2.6, and between 14.1% and
389 36.3% for RCP8.5.

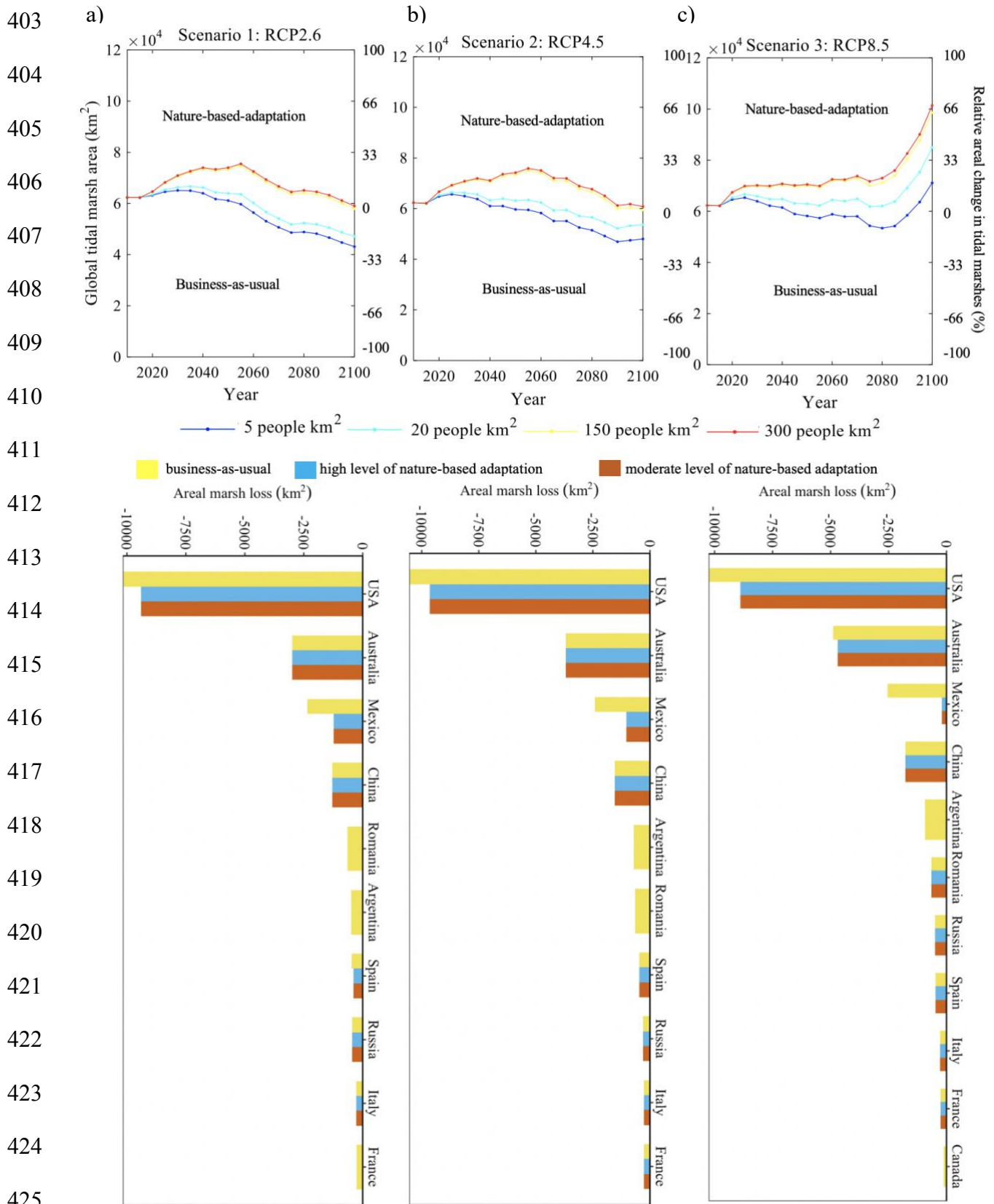


399 **Fig. 6** The relationships between tidal marsh surface elevation change and relative sea level rise under different scenarios of future sea level rise.

400 SEC is estimated by RSLR under IPCC SLR scenarios of a) RCP 2.6 as $SEC = 10^{0.001 \pm 0.0002 RSLR + 1.635 \pm 0.004} - 40$ ($R^2 = 0.33$, $p < 0.05$, F test), b) RCP4.5

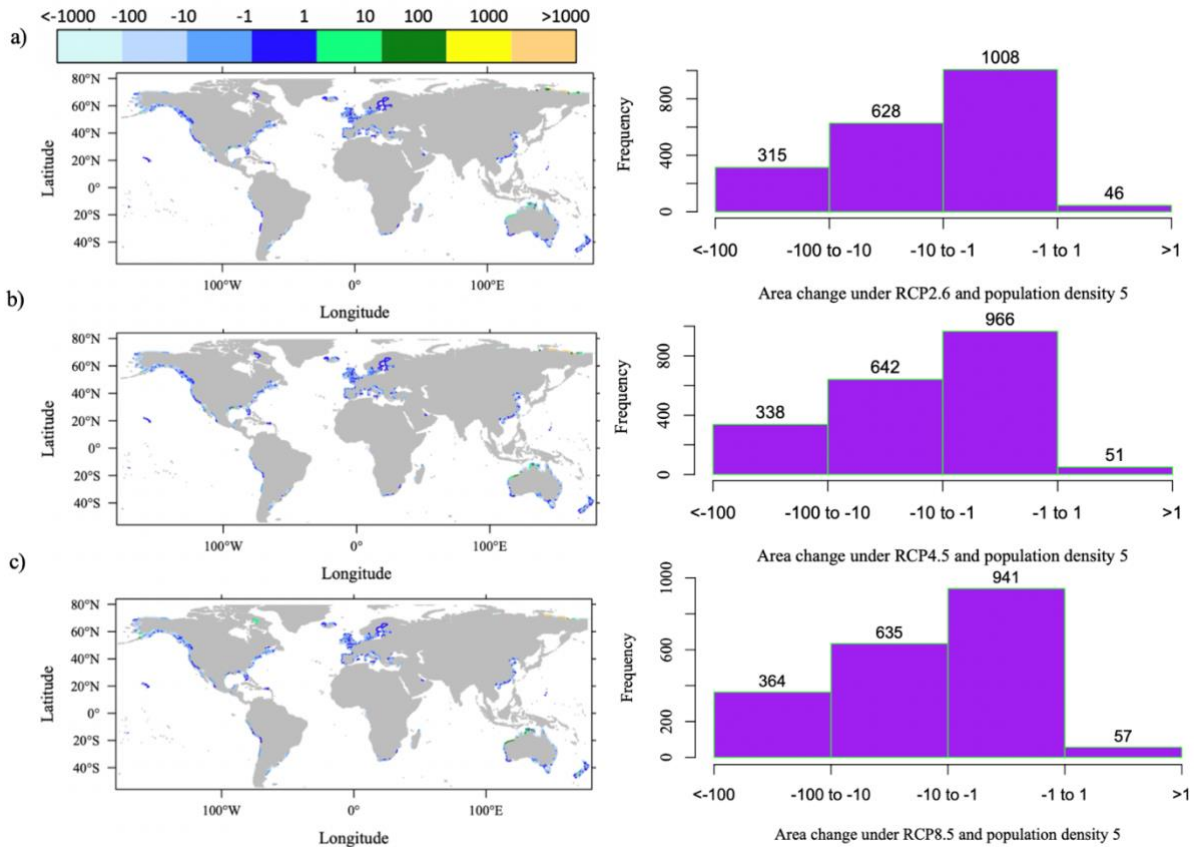
401 as $SEC = 10^{0.001 \pm 0.0002 RSLR + 1.629 \pm 0.004} - 40$ ($R^2 = 0.21$, $p < 0.05$, F test), and c) RCP 8.5 as $SEC = 10^{0.001 \pm 0.0002 RSLR + 1.627 \pm 0.004} - 40$ ($R^2 = 0.34$, $p < 0.05$, F test).

402 RSLR denotes relative sea level rise. IPCC denotes International Panel on Climate Change.



426 **Fig. 7** Changes in global tidal marsh area under different scenarios of sea level rise and
 427 thresholds of coastal population density. The projections are made on three SLR scenarios:

428 (a) RCP 2.6; (b) RCP4.5; and (c) RCP8.5. The lower panels indicate countries with hotspots
 429 of marsh areal loss ($>100 \text{ km}^2$) under different adaptation and SLR scenarios by 2100. The
 430 upper panels indicate net wetland area change estimated under different IPCC SLR scenarios.



431
 432 **Fig. 8** Frequency of absolute changes in tidal marsh area by 2100 under different scenarios of
 433 sea level rise and coastal population density of 5 people km^{-2} . Net wetland area change is
 434 estimated under different IPCC SLR scenarios of (a) RCP 2.6; (b) RCP 4.5; and (c) RCP 8.5.

435
 436 We further evaluated the resilience of tidal marshes to SLR by estimating spatially-explicit
 437 tidal marsh area changes by 2100 under different RCP scenarios (Fig. 8). At the lower
 438 boundary of the business-as-usual scenario (5 people km^{-2}), we estimated the hotspots of tidal
 439 marsh loss ($>100 \text{ km}^2$) would account for 66.8%, 67% and 68.2% of total areal marsh losses
 440 under RCP 2.6, RCP4.5 and RCP8.5, respectively. These hotspots of marsh loss mainly occur

441 in USA, China, Australia and some European countries (including Russia, Italy and France)
442 (Fig. 7).

443

444 **4 Discussion**

445 Precipitation may regulate tidal marsh SAR and SEC through sediment loadings via marine
446 or freshwater input. Sediment availability can increase during rainfall and storm events
447 (French and Spencer, 1993; Orson et al., 1998), as major storm surges can erode tidal marsh
448 edges and surfaces (Nyman et al., 1993). Tidal marsh SEC is the difference between sediment
449 deposition rate and other processes driving elevation change (Allen 1990; D'Alpaos et al.
450 2011). The influence of suspended sediments on SAR and SEC is corroborated by previous
451 modelling studies. Mineral sediment deposition is related to marsh inundation height, which
452 has a positive linear relationship with suspended sediment concentration (Temmerman et al.,
453 2003). Suspended sediment concentrations also strongly influence the maximum SLR that
454 marshes can survive (Kirwan et al., 2016). Precipitation may also have an impact on
455 freshwater input, which carries allochthonous suspended sediments to tidal marshes (McKee
456 et al., 2012). However, precipitation may not contribute significantly to sediments transported
457 to tidal marshes without fluvial input. Our results show that fluvial tidal marshes (n=317)
458 account for the highest number of sampling sites in the collated references (Fig. 5a). This
459 probably explains why precipitation is the dominant driver of SAR and SEC in our models.

460

461 Our results show that SAR predicts SEC ($R^2=0.98$), more closely than has been reported for
462 mangroves (Lovelock et al., 2015), which is consistent with an earlier finding that showed the
463 approximation of tidal marsh SAR to SEC concluded from a smaller database (Kirwan et al.,
464 2016).

465

466 Tides bring allochthonous mineral sediments and contribute to tidal marsh sediment
467 accretion, and help redistribute sediments within tidal marshes (Chmura et al., 2004). Larger
468 tidal ranges correspond to stronger tidal flow and wider intertidal regions (Rogers et al.,
469 2019), including the upper intertidal region where sediment accretion is most significant and
470 noticeable, and larger accommodation space available for increasing elevation. Further,
471 smaller tidal range means lower tidal energy, which limits inorganic sediment input from
472 marine sources, whereas macro-tides generate strong currents, which resuspend nearshore
473 sediments and transport them onto tidal marsh surfaces (Hensel et al., 1999). Thus,
474 macrotidal marshes (tidal range > 4m) are more resilient to changes in the rate of relative SLR
475 (RSLR) than microtidal marshes (tidal range < 2m)¹. In winter, the presence of sea ice affects
476 tidal inundation and thus tidal marsh SAR in high-latitude regions (Ward et al., 2014).

477

478 Autochthonous marsh production contributes to sediment accretion, changes with latitude
479 (Kirwan et al., 2009). Both tidal marsh plant shoots and roots add organic matter to the
480 sediment surface and promote mineral sediment deposition, with roots also slowing erosion
481 rates (Gedan et al., 2011). Highly productive marshes will approach a new equilibrium state
482 responding to a step change in the RSLR faster than less productive marshes (D'Alpaos et al.,
483 2011). Therefore, latitude may have an indirect impact on SAR and SEC due to the latitudinal
484 influence on marsh productivity. A similar outcome may exist also via the tidal subsidy effect
485 (Odum 1980).

486

487 Tidal flooding, geomorphology and plant traits may determine tidal marsh SAR and SEC via
488 indirect avenues, including erosion, sediment and organic/mineral matter sources. High
489 hydrodynamic energy from waves or currents during frequent flooding can result in tidal
490 marsh erosion, which causes loss of tidal marsh surfaces (Bouma et al., 2005). Tidal marsh

491 geomorphology may affect the dominant sources of sediment input, i.e. riverine or marine
492 input (Macreadie et al., 2017). Marsh plant above-ground growth can enhance the strength of
493 the feedback between tidal inundation and mineral sedimentation (Kirwan et al., 2013), while
494 below-ground components contribute mainly to organic matter accretion. Plant types with
495 different morphology and complexity may enhance the retention of sediments. However,
496 while plant traits may determine the input of organic or mineral material to sediments in tidal
497 marshes, they cannot simply account for other subsurface processes, e.g. subsidence. Shrubs
498 were reported to effectively trap sand while herbaceous plants are efficient in trapping finer
499 particles (Corenblit et al., 2009), accounting for the higher SAR with shrub occurrence. The
500 distance to creek banks may influence inputs of sediments from the water column in riverine
501 marshes (Craft et al., 1993).

502

503 Anthropogenic disturbance compromises tidal marsh SAR because the contribution of
504 resuspended sediments to tidal marsh accretion can be greatly attenuated by dredged canals,
505 spoil banks and impoundments (Boumans and Day 1994; Reed et al., 2006). Some human
506 activities (e.g. disturbance) have negative effects on SAR whereas others (e.g. restoration) do
507 not.

508

509 Our results forecast that the areal extents of tidal marshes all decline (-30.8% to -7%) in the
510 lowest SLR scenario but all increase (14.1% to 62.7%) in the highest SLR scenario under
511 different human adaptation scenarios. The areal increase is attributed to the elevation capital
512 that allows tidal marshes to migrate landward and survive through the long-term conversion
513 of terrestrial vegetation (Schuerch et al., 2019), and sediment surplus favouring vertical
514 accretion partly explained by the A/S model (Törnqvist et al., 2019). In contrast, the decrease

515 in marsh area extent in the lowest SLR scenario is mainly driven by increasing sediment
516 deficiency, which weakens the capacity of tidal marshes to keep pace with SLR.

517

518 The trend of tidal marsh areal changes is generally in line with the previous projection
519 (Schuerch et al., 2018) but occurrence of the areal peak is different. The areal peak occurs
520 around the mid-21st century under RCP 2.6 and RCP 4.5 and late-21st century under RCP8.5
521 in our projections, compared to occurrences in the late-21st century under all SLR scenarios
522 in the previous projection. The previous projection was criticized for estimating the
523 adaptation ability of coastal wetlands to SLR using a wetland adaptation score defined in
524 terms of sediment surplus or deficit, which is mistakenly represented by suspended sediment
525 concentration (Törnqvist et al., 2019). Tidal marshes are highly efficient in sequestering CO₂
526 from the atmosphere and accumulate carbon in sediments for millennial (Ouyang et al. 2014).
527 Tidal marshes, as a component of blue carbon, has been suggested for carbon abatement in
528 Australia (Kelleway et al., 2020). Our projections of their areal peak around the mid-21st
529 century are consistent with the target of a climate neutral world by 2050 as enshrined in the
530 Paris Agreement. This means that tidal marshes will contribute to the climate neutral target
531 and our study may facilitate future estimates on the contribution of tidal marshes to the target.

532

533 A previous assessment of countries with coastlines most vulnerable to SLR identified very
534 different areas of vulnerability to our modelled hotspots of marsh loss (Nicholls and
535 Cazenave, 2010). This highlights the impact of our findings - tidal marshes in countries
536 experiencing the most serious SLR will not necessarily be submerged if sufficient
537 accommodation space allows lateral migration. In contrast, at the upper boundary of the high
538 level of nature-based adaptation scenario (300 people km⁻²), the proportion of hotspots of
539 tidal marsh loss is 2.8-3.5% lower, accounting for 64-64.7% of total marsh areal losses under

540 different SLR scenarios (Supplementary Material Fig. S6). This finding highlights the
541 effectiveness of nature-based adaption for shrinking hotspots of tidal marsh loss.
542
543 Our method cannot precisely account for the fate of tidal marshes, since ecogeomorphic
544 feedbacks tend to increase rates of sediment accumulation when marshes become more
545 flooded (Kirwan et al., 2016) under future SLR scenarios. However, it is difficult to estimate
546 the extent of SAR increase due to ecogeomorphic feedbacks in the future, in particular
547 vegetation type changes at local scales (Reed et al., 2020). Moreover, geomorphologic
548 systems often react with relatively long time lags, with the response interval depending partly
549 on the magnitude, frequency and duration of energy factors (Wright and Thom, 1977). Allen
550 (1990) demonstrated a clear lag between SAR and RSLR for immature marshes. The lag was
551 analysed through a numerical model (Kirwan and Murray, 2008), and an analytical model
552 which captures the role of governing factors such as suspended sediment concentrations and
553 plant productivity. We used a 5-year lag, which may not be enough to account for feedbacks
554 at the local scale, adding uncertainty to our analysis of tidal marsh resilience to future SLR.
555 Additionally, the resolution of the satellite data in our models may have limited the precision
556 of the estimate on suspended matter, precipitation and sea ice. However, our mathematical
557 model provides new insights on assessing the resilience of tidal marshes to SLR with real
558 elevation change data by incorporating both vertical sediment accretion and lateral landward
559 migration.

560

561 **Data availability**

562 The satellite data that support the findings of this study are downloaded from

563 <http://hermes.acri.fr/index.php?class=archive> (total suspended matter) and

564 <https://www.metoffice.gov.uk/hadobs/hadisst/> (sea ice),

565 <https://www.esrl.noaa.gov/psd/data/gridded/data.gpcc.html> and
566 <http://catalogue.ceda.ac.uk/uuid/3f8944800cc48e1cbc29a5ee12d8542d> (precipitation).
567 Satellite and literature data are available in ‘figshare’ with the identifier doi:
568 10.6084/m9.figshare.7545212 or upon reasonable request sent to the corresponding author.
569 Computer code is available from Xiaoguang Ouyang upon reasonable request.

570

571 **Supplementary Material**

572 Fig. S1

573 Fig. S2

574 Fig. S3

575 Fig. S4

576 Fig. S5

577 Fig. S6

578 Table S1

579

580 **References**

581 Allen, J. (1990). Constraints on measurement of sea-level movements from salt-marsh
582 accretion rates. *Journal of the Geological Society*, 147, 5-7.

583 Allen, J. (2000). Morphodynamics of Holocene salt marshes: a review sketch from the
584 Atlantic and Southern North Sea coasts of Europe. *Quaternary Science Reviews*, 19,
585 1155-1231.

586 Andersen, T., Svinth, S., & Pejrup, M. (2011). Temporal variation of accumulation rates on a
587 natural salt marsh in the 20th century—The impact of sea level rise and increased
588 inundation frequency. *Marine Geology*, 279, 178-187.

589 Anisfeld, S. C., Hill, T. D., & Cahoon, D. R. (2016). Elevation dynamics in a restored versus
590 a submerging salt marsh in Long Island Sound. *Estuarine, Coastal and Shelf Science*,
591 *170*, 145-154.

592 Barbier, E. B. A spatial model of coastal ecosystem services. *Ecol. Econ*, **78**, 70-79 (2012).

593 Bivand, R. and Rundel, C. (2020). rgeos: Interface to Geometry Engine - Open Source
594 ('GEOS'). R package version 0.5-3. <https://CRAN.R-project.org/package=rgeos>

595 Blondeau-Patissier, D., Schroeder, T., Brando, V., Maier, S., Dekker, A., & Phinn, S. (2014).
596 ESA-MERIS 10-year mission reveals contrasting phytoplankton bloom dynamics in
597 two tropical regions of Northern Australia. *Remote Sensing*, *6*, 2963-2988.

598 Bouma, T., Vries, M. D., Low, E., Kusters, L., Herman, P., Tanczos, I., Temmerman, S.,
599 Hesselink, A., Meire, P., & Van Regenmortel, S. (2005). Flow hydrodynamics on a
600 mudflat and in salt marsh vegetation: identifying general relationships for habitat
601 characterisations. *Hydrobiologia*, *540*, 259-274.

602 Boumans, R. M., & Day, J. W. (1994). Effects of two Louisiana marsh management plans on
603 water and materials flux and short-term sedimentation. *Wetlands*, *14*, 247-261.

604 Breithaupt, J. L., Smoak, J. M., Byrne, R. H., Waters, M. N., Moyer, R. P., & Sanders, C. J.
605 (2018). Avoiding timescale bias in assessments of coastal wetland vertical change.
606 *Limnology and Oceanography*, *63*, 477-495.

607 Cahoon, D. R., Reed, D. J., & Day, J. W. (1995). Estimating shallow subsidence in microtidal
608 salt marshes of the southeastern United States: Kaye and Barghoorn revisited. *Marine*
609 *Geology*, *128*, 1-9.

610 Cahoon, D. R., Perez, B. C., Segura, B. D., & Lynch, J. C. (2011). Elevation trends and
611 shrink–swell response of wetland soils to flooding and drying. *Estuarine, Coastal and*
612 *Shelf Science*, *91*, 463-474.

613 Chmura, G. L., & Hung, G. A. (2004). Controls on salt marsh accretion: a test in salt marshes
614 of Eastern Canada. *Estuaries*, *27*, 70-81.

615 Corenblit, D., Steiger, J., Gurnell, A. M., Tabacchi, E., & Roques, L. (2009). Control of
616 sediment dynamics by vegetation as a key function driving biogeomorphic succession
617 within fluvial corridors. *Earth Surface Processes and Landforms*, *34*, 1790-1810.

618 Costanza, R. *et al.* Changes in the global value of ecosystem services. *Global Environ.*
619 *Change* **26**, 152-158 (2014).

620 Craft, C., Seneca, E. D., & Broome, S. W. (1993). Vertical accretion in microtidal regularly
621 and irregularly flooded estuarine marshes. *Estuarine, Coastal and Shelf Science*, *37*,
622 371-386.

623 Craft, C. (2007). Freshwater input structures soil properties, vertical accretion, and nutrient
624 accumulation of Georgia and US tidal marshes. *Limnology and Oceanography*, *52*,
625 1220-1230.

626 Craft, C., Clough, J., Ehman, J., Joye, S., Park, R., Pennings, S., Guo, H., & Machmuller, M.
627 (2008). Forecasting the effects of accelerated sea-level rise on tidal marsh ecosystem
628 services. *Frontiers in Ecology and the Environment*, *7*, 73-78.

629 D'Alpaos, A., Mudd, S. M., & Carniello, L. (2011). Dynamic response of marshes to
630 perturbations in suspended sediment concentrations and rates of relative sea level rise.
631 *Journal of Geophysical Research: Earth Surface*, *116*(F4).

632 Dijkstra, L., & Poelman, H. (2014). A Harmonised Definition of Cities and Rural Areas: the
633 New Degree of Urbanisation, European Commission Working Paper WP.

634 Doody, J. P. (2004). 'Coastal squeeze'—an historical perspective. *Journal of Coastal*
635 *Conservation*, *10*, 129-138.

636 French, J. R., & Spencer, T. (1993). Dynamics of sedimentation in a tide-dominated
637 backbarrier salt marsh, Norfolk, UK. *Marine Geology*, *110*, 315-331.

638 Friedrichs, C. T., & Perry, J. E. (2001). Tidal salt marsh morphodynamics: a synthesis.
639 *Journal of Coastal Research*, Special issue 27, 7-37.

640 Gabler, R. E., Petersen, J. F., Trapasso, L. M., Sack, D., Sager, R. J., & Wise, D. L. (2007).
641 *Essentials of Physical Geography, Eighth Edition*. CA, USA: Thomson Higher
642 Education.

643 Gedan, K. B., Kirwan, M. L., Wolanski, E., Barbier, E. B., & Silliman, B. R. (2011). The
644 present and future role of coastal wetland vegetation in protecting shorelines:
645 answering recent challenges to the paradigm. *Climatic Change*, 106, 7-29.

646 Gedan, K. B., Silliman, B. R., & Bertness, M. D. (2009). Centuries of human-driven change
647 in salt marsh ecosystems. *Annual Review of Marine Science*, 1, 117-141.

648 Greenwell, B., B. Boehmke, J. Cunningham, & GBM, D. (2019). gbm: Generalized Boosted
649 Regression Models. R package version 2.1.5. [https://CRAN.R-](https://CRAN.R-project.org/package=gbm)
650 [project.org/package=gbm](https://CRAN.R-project.org/package=gbm).

651 Hensel, P. E., Day Jr, J. W., & Pont, D. (1999). Wetland vertical accretion and soil elevation
652 change in the Rhone River delta, France: the importance of riverine flooding. *Journal*
653 *of Coastal Research*, 15, 668-681.

654 Hijmans, R. J., S. Phillips, J. Leathwick, & Elith., J. (2017). dismo: Species Distribution
655 Modeling. R package version 1.1-4. <https://CRAN.R-project.org/package=dismo>.

656 Hinkel, J., Lincke, D., Vafeidis, A. T., Perrette, M., Nicholls, R. J., Tol, R. S., Marzeion, B.,
657 Fettweis, X., Ionescu, C., & Levermann, A. (2014). Coastal flood damage and
658 adaptation costs under 21st century sea-level rise. *Proceedings of the National*
659 *Academy of Sciences USA*, 111, 3292-3297.

660 IPCC. (2014). Climate Change 2014: Synthesis Report. Contribution of Working Groups I, II
661 and III to the Fifth Assessment Report of the Intergovernmental Panel on Climate

662 Change [Core Writing Team, R.K. Pachauri and L.A. Meyer (eds.)]. IPCC, Geneva,
663 Switzerland, 151 pp.

664 Kelleway, J.J., Serrano, O., Baldock, J.A., Burgess, R., Cannard, T., Lavery, P.S., Lovelock,
665 C.E., Macreadie, P.I., Masqué, P., Newnham, M. and Saintilan, N. (2020) A national
666 approach to greenhouse gas abatement through blue carbon management. *Global*
667 *Environmental Change*, 63, 102083.

668 Kelley, J. T., Belknap, D. F., Jacobson Jr, G. L., & Jacobson, H. A. (1988). The morphology
669 and origin of salt marshes along the glaciated coastline of Maine, USA. *Journal of*
670 *Coastal Research*, 4, 649-665.

671 Kirwan, M. L., & Murray, A. B. (2008). Tidal marshes as disequilibrium landscapes? Lags
672 between morphology and Holocene sea level change. *Geophysical Research Letters*
673 35, L24401.

674 Kirwan, M. L., Guntenspergen, G. R., & Morris, J. T. (2009). Latitudinal trends in *Spartina*
675 *alterniflora* productivity and the response of coastal marshes to global change. *Global*
676 *Change Biology*, 15, 1982-1989.

677 Kirwan, M. L., & Megonigal, J. P. (2013). Tidal wetland stability in the face of human
678 impacts and sea-level rise. *Nature*, 504, 53-60.

679 Kirwan, M. L., Temmerman, S., Skeeahan, E. E., Guntenspergen, G. R., & Fagherazzi, S.
680 (2016). Overestimation of marsh vulnerability to sea level rise. *Nature Climate*
681 *Change*, 6, 253-260.

682 Koch, E. W., Barbier, E. B., Silliman, B. R., Reed, D. J., Perillo, G. M. E., Hacker, S. D.,
683 Granek, E. F., Primavera, J. H., Muthiga, N., Polasky, S., Halpern, B. S., Kennedy, C.
684 J., Kappel, C. V., & Wolanski, E. (2009). Non-linearity in ecosystem services:
685 temporal and spatial variability in coastal protection. *Frontiers in Ecology and the*
686 *Environment*, 7, 29-37.

687 Kolker, A. S., Goodbred, S. L., Hameed, S., & Cochran, J. K. (2009). High-resolution records
688 of the response of coastal wetland systems to long-term and short-term sea-level
689 variability. *Estuarine, Coastal and Shelf Science*, 84, 493-508.

690 Leonardi, N., Carnacina, I., Donatelli, C., Ganju, N. K., Plater, A. J., Schuerch, M., &
691 Temmerman, S. (2018). Dynamic interactions between coastal storms and salt
692 marshes: A review. *Geomorphology*, 301, 92-107.

693 Lovelock, C. E., Cahoon, D. R., Friess, D. A., Guntenspergen, G. R., Krauss, K. W., Reef, R.,
694 Rogers, K., Saunders, M. L., Sidik, F., & Swales, A. (2015). The vulnerability of
695 Indo-Pacific mangrove forests to sea-level rise. *Nature*, 526, 559-563.
696 doi:10.1038/nature15538.

697 Ma, Z., Melville, D. S., Liu, J., Chen, Y., Yang, H., Ren, W., Zhang, Z., Piersma, T., & Li, B.
698 (2014). Rethinking China's new great wall. *Science*, 346, 912-914.

699 Macreadie PI, O. Q., Kelleway JJ, Serrano O, Carnell PE, Ewers Lewis CJ, Atwood TB,
700 Sanderman J, Baldock J, Connolly RM, Duarte CM, Lavery PS, Steven A, Lovelock
701 CE. . (2017). Carbon sequestration by Australian tidal marshes. *Scientific Reports*, 7,
702 44071.

703 Marani, M., Belluco, E., Ferrari, S., Silvestri, S., D'Alpaos, A., Lanzoni, S., Feola, A., &
704 Rinaldo, A. (2006). Analysis, synthesis and modelling of high-resolution observations
705 of salt-marsh eco-geomorphological patterns in the Venice lagoon. *Estuarine, Coastal
706 and Shelf Science*, 69, 414-426.

707 Marani, M., d'Alpaos, A., Lanzoni, S., & Santalucia, M. (2011). Understanding and
708 predicting wave erosion of marsh edges. *Geophysical Research Letters*, 38, L21401.

709 McKee, K., Rogers, K., & Saintilan, N. (2012). Response of salt marsh and mangrove
710 wetlands to changes in atmospheric CO₂, climate, and sea level. In *Global change and
711 the function and distribution of wetlands* (pp. 63-96): Springer.

712 Mittermeier, R. A., Mittermeier, C. G., Brooks, T. M., Pilgrim, J. D., Konstant, W. R., da
713 Fonseca, G. A., & Kormos, C. (2003). Wilderness and biodiversity conservation.
714 *Proceedings of the National Academy of Sciences USA*, *100*, 10309-10313.

715 Nicholls, R. J., & Cazenave, A. (2010). A Sea-level rise and its impact on coastal zones.
716 *Science* *328*, 1517-1520.

717 Nyman, J. A., DeLaune, R., Roberts, H., & Patrick Jr, W. (1993). Relationship between
718 vegetation and soil formation in a rapidly submerging coastal marsh. *Marine Ecology*
719 *Progress Series*, *96*, 269-279.

720 Odum, E. P. (1980). The status of three ecosystem-level hypotheses regarding salt marsh
721 estuaries: tidal subsidy, outwelling, and detritus-based food chains. In *Estuarine*
722 *perspectives* (pp. 485-495): Elsevier.

723 Orson, R., Warren, R., & Niering, W. (1998). Interpreting sea level rise and rates of vertical
724 marsh accretion in a southern New England tidal salt marsh. *Estuarine, Coastal and*
725 *Shelf Science*, *47*, 419-429.

726 Ouyang, X., & Lee, S. Y. (2014). Updated estimates of carbon accumulation rates in coastal
727 marsh sediments. *Biogeosciences*, *11*, 5057-5071. doi:10.5194/bg-11-5057-2014.

728 Ouyang, X., Lee, S. Y., & Connolly, R. M. (2017). The role of root decomposition in global
729 mangrove and saltmarsh carbon budgets. *Earth-Science Reviews*, *166*, 53-63.
730 doi:10.1016/j.compchemeng.2016.09.009.

731 Parkinson, R. W., Craft, C., DeLaune, R. D., Donoghue, J. F., Kearney, M., Meeder, J. F.,
732 Morris, J., & Turner, R. E. (2017). Marsh vulnerability to sea-level rise. *Nature*
733 *Climate Change*, *7*, 756.

734 Pierce, D. (2017). ncdf4: Interface to Unidata netCDF (Version 4 or Earlier) Format Data
735 Files. R package version 1.16. <https://CRAN.R-project.org/package=ncdf4>.

736 R Core Team. (2014). A Language and Environment for Statistical Computing, R Foundation
737 for Statistical Computing, Vienna, Austria.

738 Rayner, N., Parker, D. E., Horton, E., Folland, C., Alexander, L., Rowell, D., Kent, E., &
739 Kaplan, A. (2003). Global analyses of sea surface temperature, sea ice, and night
740 marine air temperature since the late nineteenth century. *Journal of Geophysical*
741 *Research: Atmospheres*, 108(D14), 4407. doi:10.1029/2002JD002670.

742 Reed, D. J., Peterson, M. S., & Lezina, B. J. (2006). Reducing the effects of dredged material
743 levees on coastal marsh function: sediment deposition and nekton utilization.
744 *Environmental Management*, 37, 671-685.

745 Reed, D., Wang, Y., Meselhe, E., & White, E. (2020). Modeling wetland transitions and loss
746 in coastal Louisiana under scenarios of future relative sea-level
747 rise. *Geomorphology*, 352, 106991.

748 Rogers, K., Wilton, K., & Saintilan, N. (2006). Vegetation change and surface elevation
749 dynamics in estuarine wetlands of southeast Australia. *Estuarine, Coastal and Shelf*
750 *Science*, 66, 559-569.

751 Rogers, K., Kelleway, J., Saintilan, N., Megonigal, J., Adams, J., Holmquist, J., Lu, M.,
752 Schile-Beers, L., Zawadzki, A., Mazumder, D., & Woodroffe, C. (2019). Wetland
753 carbon storage controlled by millennial-scale variation in relative sea-level rise.
754 *Nature*, 567, 91-95. doi:10.1038/s41586-019-0951-7.

755 Rybczyk, J. M., & Cahoon, D. R. (2002). Estimating the potential for submergence for two
756 wetlands in the Mississippi River Delta. *Estuaries*, 25, 985-998.

757 Schuerch, M., Spencer, T., Temmerman, S., Kirwan, M. L., Wolff, C., Lincke, D., McOwen,
758 C. J., Pickering, M. D., Reef, R., & Vafeidis, A. T. (2018). Future response of global
759 coastal wetlands to sea-level rise. *Nature*, 561, 231-234.

760 Schuerch, M., Spencer, T., Temmerman, S., Kirwan, M. L., Wolff, C., Lincke, D., McOwen,
761 C. J., Pickering, M. D., Reef, R., Vafeidis, A. T., Jochen Hinkel, & Robert J. Nicholls,
762 S. B. (2019). Reply to “Global coastal wetland expansion under accelerated sea-level
763 rise is unlikely”, EarthArXiv. doi:10.31223/osf.io/d2nhs).

764 Temmerman, S., Govers, G., Meire, P., & Wartel, S. (2003). Modelling long-term tidal marsh
765 growth under changing tidal conditions and suspended sediment concentrations,
766 Scheldt estuary, Belgium. *Marine Geology*, 193, 151-169.

767 Thom, R. M. (1992). Accretion rates of low intertidal salt marshes in the Pacific Northwest.
768 *Wetlands*, 12, 147-156.

769 Titus, J., Wang, J. (2008) Background Documents Supporting Climate Change Science
770 Program synthesis and Assessment Product 4.1 Coastal Elevations and Sensitivity to
771 Sea-Level Rise. EPA 430R07004. United States Environmental Protection Agency,
772 Washington, DC. 354 pp.

773 Törnqvist, T., Cahoon, D., Day, J., & Morris, J. (2019). Global coastal wetland expansion
774 under accelerated sea-level rise is unlikely. EarthArXiv. July 16.
775 doi:10.31223/osf.io/d2nhs.

776 Ward, R. D., Teasdale, P. A., Burnside, N. G., Joyce, C. B., & Sepp, K. (2014). Recent rates
777 of sedimentation on irregularly flooded Boreal Baltic coastal wetlands: responses to
778 recent changes in sea level. *Geomorphology*, 217, 61-72.

779 Wolters, M., Garbutt, A., Bekker, R. M., Bakker, J. P., & Carey, P. D. (2008). Restoration of
780 salt-marsh vegetation in relation to site suitability, species pool and dispersal traits.
781 *Journal of Applied Ecology*, 45, 904-912.

782 Wright, L., & Thom, B. (1977). Coastal depositional landforms: a morphodynamic approach.
783 *Progress in Physical Geography*, 1, 412-459.

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Declaration of interests

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The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:



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