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Multi-temporal relative landslide risk analysis for sustainable development of rapidly growing cities

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Abstract

In the last decades, developing countries have experienced an increase in impact of natural disasters due to both the ongoing climate change and the sustained expansion of urban areas. Intrinsic vulnerability of settlements due to poverty and poor governance, as well as the lack of tools for urban occupation planning and mitigation protocols, have made such impact particularly severe. Cuenca (Ecuador) is a significant example of a city that in the last decades has experienced considerable population growth and an associated increasing of loss due to landslide occurrence. Despite such effects, updated urban planning tools are absent, a condition that suggested an evaluation of multi-temporal relative landslide risk, here presented based on updated data depicting the spatial distribution of landslides and their predisposing factors, as well as population change between 2010 and 2020. In addition, a multi-temporal analysis accounting for risk change between 2010 and 2020 has been carried out. Due to the absence of spatially distributed data about the population, electricity supply contract data have been used as a proxy of the population. Results indicate that current higher relative risk is estimated for municipalities (*parroquias*) located at the southern sector of the study area (i.e. *Turi, Valle, Santa Ana, Tarqui* and *Paccha*). Moreover, the multi-temporal analysis indicates that most municipalities of the city located in the hilly areas that bound the center (i.e. *Sayausi, San*

46 *Joaquin, Tarqui, Valle, Sidcay, Banos, Sidcay, Ricaurte, Paccha and Chiquintad*), experiencing
47 sustained population growth, will be exposed to an increased risk with a consistently growing trend.
48 This information is consistent with landslide susceptibility data derived by a machine learning-based
49 analysis that indicate higher susceptibility to landslides in hilly areas surrounding the city center. The
50 obtained relative risk maps can be considered as a useful tool for guiding land-planning, occupation
51 restriction and early warning strategy adoption. The used methodological approach, accounting for
52 landslide susceptibility and population variation through proxy data analysis, has the potential to be
53 applied in a similar context of growing-population cities of low to mid-income countries, where data,
54 usually needed for a comprehensive landslide risk analysis, are only partly available.

55

56 **Keywords:** Landslide susceptibility; Machine learning algorithm; Relative risk assessment; Cuenca;
57 Ecuador; Latin America.

58

59 **1. Introduction**

60 In the last decades, the ongoing climate change, associated with global population growth and the
61 related expansion of urbanized areas, has been responsible for an increase in frequency and impact
62 of natural disasters due to floods, landslides and wildfires (Knox 1993; Xu et al. 2013; Altan et al.
63 2015; Arnell and Gosling 2016; Gariano and Guzzetti 2016; Di Napoli et al. 2020a). Developing
64 countries have experienced an even much severe impact, because people are often concentrated in
65 high-hazard urban areas where housing is highly vulnerable due to poor building, and early warning
66 systems are commonly absent (Zorn 2018; Aguirre-Ayerbe et al. 2020). The dimension of such
67 impact can be easily understood considering that between 1996 and 2015 approximately 90% of
68 disaster-related deaths occurred in mid to low-income countries
69 (<http://reliefweb.int/report/world/poverty-death-disaster-and-mortality-1996-2015>).

70 Vulnerability factors such as poverty, poor governance, and the lack of experience in facing natural
71 disaster are responsible for this effect disproportion (i.e., deaths concentration in developing

72 countries; Petley, 2012). The lack of tools for urban occupation planning, prescriptions definition,
73 and mitigation protocols, is a further element of vulnerability that particularly applies to rapidly
74 growing urban areas prone to floods and landslides. For this reason, an evaluation and prevention of
75 exposure to geohazards, in terms of susceptibility and hazard, is a fundamental step for the correct
76 environment planning and management, as shown by several scientific contributions in this field
77 (Goetz et al. 2015; Guerriero et al. 2018, 2020a, b; Lombardo et al. 2020; Segoni et al. 2020; Di
78 Napoli et al. 2021; Novellino et al. 2021; Allocca et al. 2021).

79 In Latin-American countries, between 2004 and 2013, 611 landslides triggered by rainfall and
80 earthquakes have been responsible for approximately 12000 deaths (Sepúlveda and Petley 2015). A
81 relevant example is an event that, in 2017, involved the city of Mocoa in southern Colombia, which
82 killed more than 300 people and destroyed 130 houses (García-Delgado et al. 2019). While the spatial
83 distribution of such events is consistently related to a combination of slope morphometry, rainfall
84 distribution and population density, poverty is a controlling factor of the impact to people particularly
85 relevant in urban areas. Indeed, the presence of informal settlements and their localization has a big
86 impact on the number of fatalities. In such conditions, landslide susceptibility and risk maps represent
87 useful tools to develop land-planning strategies for preventing such kinds of impact and supporting
88 the sustainable development of cities (Musakwa and van Niekerk 2015; AlQahtany and Abubakar
89 2020).

90 Landslide susceptibility indicates the probability of a slope failure occurring in an area depending on
91 its geomorphological peculiarities (van Westen et al. 2003; Guzzetti et al. 2006; Reichenbach et al.
92 2018). It differs from hazard, since does not directly consider any evaluation of the expected
93 magnitude of an event and its recurrence time (Fell et al. 2008). A landslide susceptibility map
94 spatially reproduces the landslide occurrence likelihood providing an overview of areas that need
95 prescriptions in settlement development perspective (Chen et al. 2020; Di Napoli et al. 2020b; Zhang
96 et al. 2020; Arabameri et al. 2021). Landslide risk depends on the characteristics of elements at risk,
97 their vulnerability and the temporal-spatial probability of occurrence of a damaging landslide event.

98 Risk maps are powerful tools since they consider also the characteristics of exposed elements
99 providing potential damage scenarios (Bignami et al. 2018; Novellino et al. 2021). In general terms,
100 landslide risk is evaluated through a multi-step analysis including *i*) hazard identification, *ii*) hazard
101 assessment, *iii*) inventory of elements at risk and exposure, *iv*) vulnerability assessment and *v*) risk
102 estimation (Dai et al. 2002; Glade et al. 2006; van Westen et al. 2008; Corominas et al. 2014). Due
103 to the frequent lack of landslide occurrence timing data, risk is often evaluated by adopting a
104 simplified approach based upon susceptibility scenarios rather than hazard (Ercanoglu 2008; Fell et
105 al. 2008; Arabameri et al. 2017). An alternative hazard can be estimated on the basis of the return
106 period of landslide triggering events (Grelle et al. 2014). A further element of simplification generally
107 relates to vulnerability estimation that, especially for heterogenous settlements, could be very
108 challenging to be correctly estimated and might cause diffuse under-or over-estimation of landslide
109 risk (Glade 2003; Li et al. 2010; Mavrouli et al. 2014). The concept of relative risk well applies to
110 regions where a reduced complexity estimation has to be preferred due to limits in data availability,
111 such as in rapidly growing urban areas of developing countries (Andrejev et al. 2017).

112 The city of Cuenca (Ecuador) is an example of a rapidly growing city that in the last decades has
113 experienced a sustained increase in population, having reached 400000 units. Magnitude of the
114 change can be estimated considering that thirty years ago the total population was around 200000
115 units. Due to this growth, the urban area has consistently expanded from its original position within
116 the *Tomebamba* River floodplain occupying surrounding hilly slopes. Such slopes are prone to
117 landslides, so that damage to settlements and infrastructures are very frequent preventing a
118 sustainable and safe development of the city (Miele et al. 2021). On this basis, an analysis of the
119 relative risk to landslide and its multi-temporal variation as a function of population growth in the
120 city of Cuenca was completed with the aim of providing *i*) an updated overview of relative risk
121 exposure to landslides of population and *ii*) a general tool for supporting land planning in rapidly
122 growing cities suffering the effect of natural hazards such due to the lack of planning instruments.
123 Indeed, for the city of Cuenca, a first attempt of natural hazards assessment dates back to the early

124 1990s with a pilot project called PRECUPA (PREvention ECuador CUenca PAute). In the
125 perspective of the analysis, an integrated method consisting of *i*) landslide inventory construction
126 through both remote-sensing data analysis and field observations, *ii*) machine-learning-based
127 susceptibility assessment by using Maximum Entropy algorithm, *iii*) electricity supply contract
128 analysis for exposure quantification and *iv*) relative risk estimation, was used.

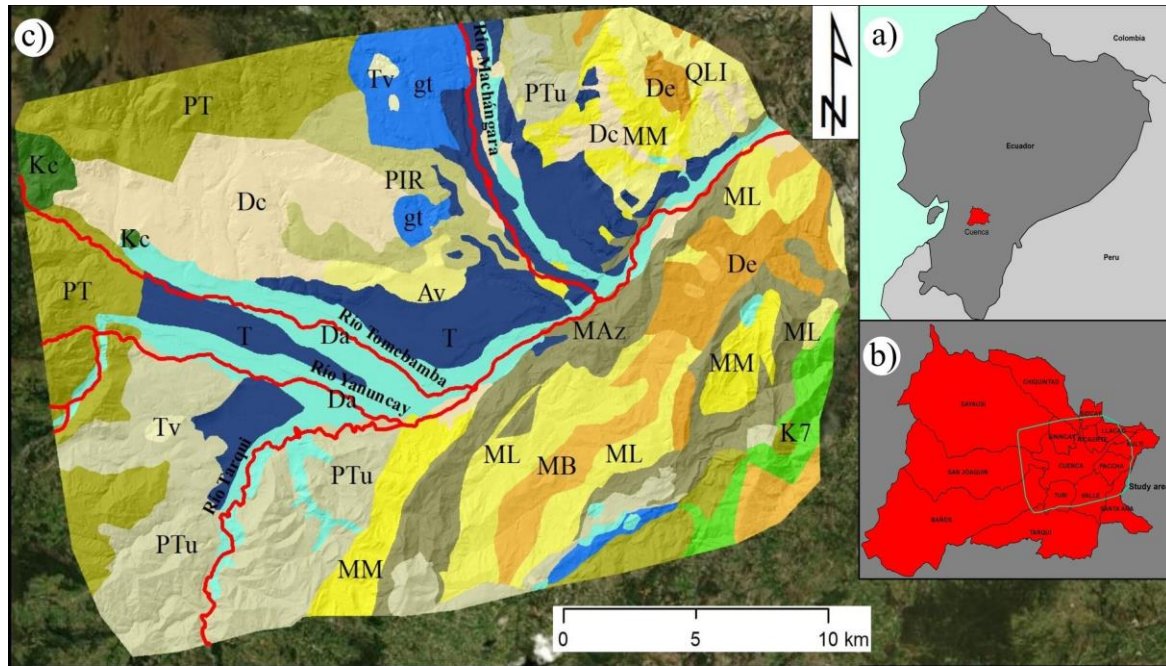
129

130 **2. Study area**

131 The study area comprises the city of Cuenca and the surrounding hilly area being increasingly
132 occupied by settlements (Fig. 1). Cuenca is the capital of the *Azuay* province of Ecuador and extends
133 over an area of approximately 124 km². In 1999, the historic center of the city has been inserted in
134 the UNESCO World Heritage Site list due to its importance as a cultural and governmental center of
135 the Canari and Inca civilizations, and for being an example of renaissance urban planning in the
136 Americas during the Spanish colonial period.

137 The city lies within an inter-Andean valley, which was formed following a compressional
138 deformation controlled by major NE-trending faults (Noblet et al. 1988; Hungerbühler et al. 2002).
139 The geology of this area is represented by Mesozoic marine and subaerial sedimentary deposits,
140 covering the Paleozoic metamorphic basement (Noblet et al. 1988). The sedimentary series is more
141 than 2400 and 3500 m thick and is formed by two main sequences separated by a regional
142 unconformity. The lower sequence consists of fluvial and brackish delta plain deposits containing
143 ubiquitous metamorphic pebbles from the Cordillera Real. From the bottom, this series is made up of
144 the *Biblián*, *Loyola*, *Azogues* and *Mangan* Formations that include sandy clays, laminated shales with
145 gypsum, tuffaceous sandstones, siltstones and conglomeratic sandstones. The upper sedimentary
146 sequence is composed of volcanic clast-bearing rocks of the *Turi* Formation, which is divided into
147 *Turi* and *Santa Rosa* members, and consists of tuffaceous coarse sandstones, volcanic clast-supported
148 conglomerates, matrix-supported volcanic breccias and minor tuff layers. Furthermore, the late
149 Miocene volcanic series of *Tarqui* Formation crops out in the area unconformably covering a wide

150 range of volcanic and Tertiary sedimentary formations. The *Tarqui* formation is formed by two
 151 members: *i*) the *Tarqui* Member formed by poorly consolidated and intensely weathered red volcanic
 152 airfall deposits and *ii*) the *Llacao* Member represented by debris-flow deposits of volcanoclastic
 153 materials (Steinmann 1997).



154
 155 Fig. 1. a) Geographic location of the study area; b) Cuenca municipalities, the green box shows the location of the
 156 considered area for the analysis; c) Cuenca geological map (refer to Table 1 for geology types' code description). Red
 157 lines represent the four principal rivers that cross Cuenca town.

158 This area is known for frequent landslides involving settlements and infrastructures (Fig. 2). For
 159 instance, the buildings of the Faculty of Philosophy of the University of *Azuay* is consistently affected
 160 by slow-moving landslides and periodically damaged by slope deformation (Sellers et al., 2020). On
 161 March 29, 1993, a large landslide (20 million m³) took place northeast of Cuenca city, damming the
 162 *Paute* river and causing the formation of an artificial lake that flooded fertile land and destroying
 163 houses, roads, railways and a regional thermoelectric plant (Plaza et al. 2011). In addition, the
 164 segment of the Pan-American Highway crossing the city is continuously affected by landslides
 165 inducing damage and, in some cases, vehicle accidents (Miele et al. 2021). The high frequency of
 166 landslides is related to the significant yearly rainfall amount (around 900 mm), the extremely high
 167 frequency of earthquakes of significant magnitude (ranging from 4.0 to 4.9 Mw according to

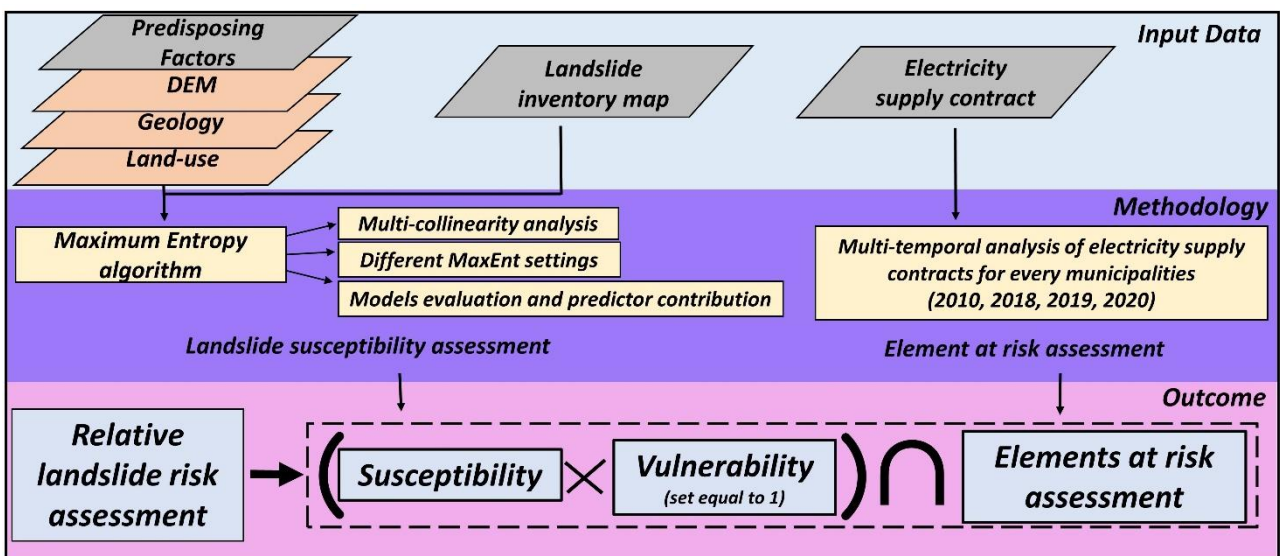
168 Ecuador's Geophysical Institute, <https://www.igepn.edu.ec/>) and the predisposing action of slope
 169 morphology and geological characteristics.



171 Fig. 2. a) House completely destroyed by a mass movement; b) house partially damaged due to a landslide; c) an
 172 example of low-risk perception by the local population. The structure was built close to a sub-vertical rock wall which
 173 is very prone to mass movements such as rockfalls, topples and slides (photo: M. Ramondini).

174 3. Data and methods

175 To evaluate the relative risk to landslides of the population of the city of Cuenca and its multi-
 176 temporal variation in relation to population growth, a method consisting of landslide mapping,
 177 machine-learning-based susceptibility analysis, population growth estimation through energy supply
 178 contract analysis and relative risk evaluation was executed (Fig. 3). Below, details of data and
 179 methods used for the estimation are provided.



181
 182 Fig. 3. Flowchart showing the methodology for relative landslide risk assessment.
 183

184 **3.1. Landslide inventory map**

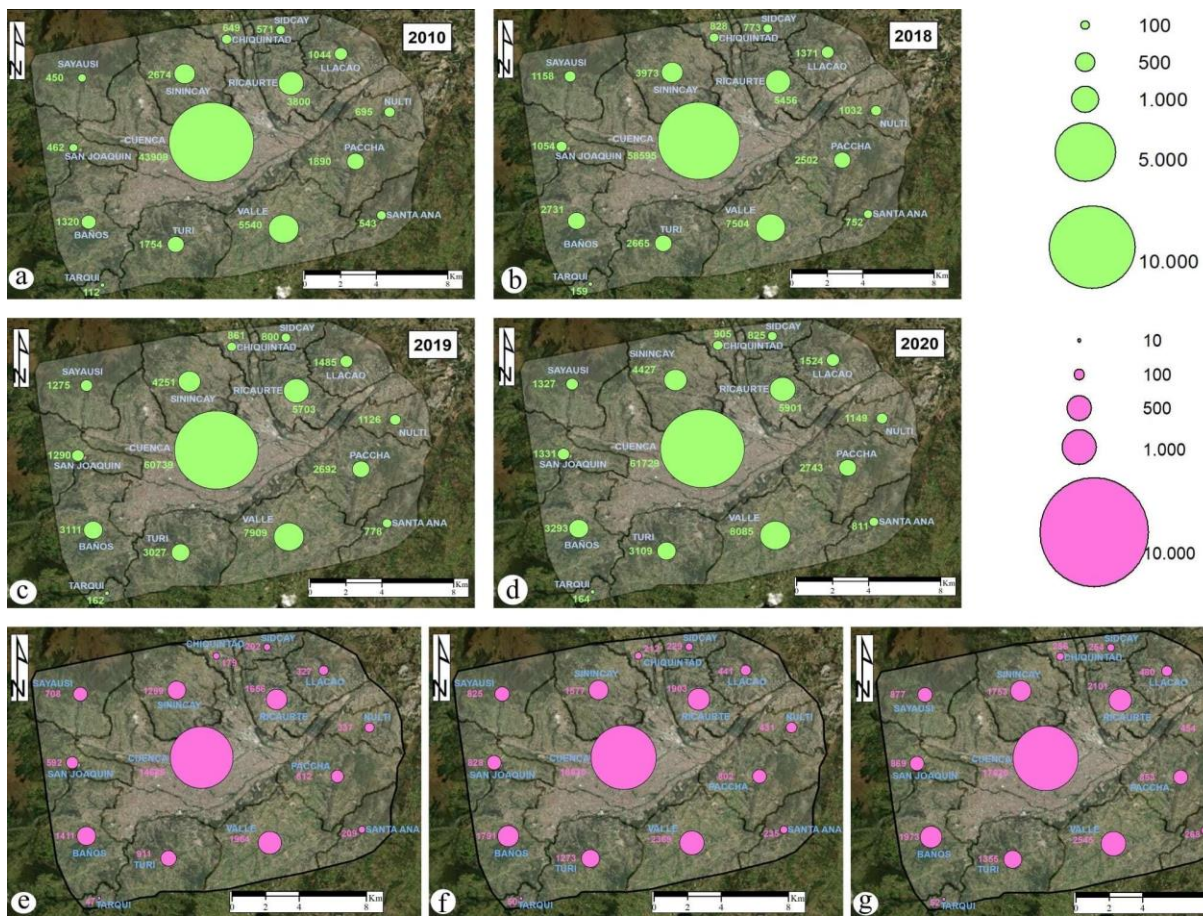
185 In the perspective of evaluating landslide relative risk to people, the assessment of susceptibility to
186 landslides was carried out using the available landslide inventory prepared by Miele et al. (2021) for
187 the area surrounding the Pan-American Highway integrated by further analyses to extend the
188 inventory to the study area. Following the method used by Miele et al. (2021), a landslide inventory
189 was derived by interferometric data, visual interpretation of aerial imagery and field surveys.
190 Sentinel-1A and B images (ascending and descending pass) acquired between October 2016 and May
191 2019 were processed using the Coherent Pixel Technique - Temporal Phase Coherence (CPT-TPC)
192 approach (Mora et al. 2003; Iglesias et al. 2015). The obtained Line of Sight mean displacement rates
193 were post-processed through the application of the kernel density estimation (KDE) algorithm,
194 allowing to identify unstable areas (UAs) affected by ongoing deformations (Di Martire et al. 2016;
195 Guerriero et al. 2019; Ammirati et al. 2020). Identified unstable areas were used as a guide for aerial
196 imagery interpretation and subsequent field surveys. The latter were carried out between September
197 2020 and March 2021 using 1:5000 topographical maps as a basemap. Landslide areas derived by
198 image analysis and field surveys were digitized into a GIS environment and classified according to
199 Cruden and Varnes (1996).

200

201 **3.2. Electricity supply contracts analysis**

202 The relative landslide risk evaluation has been completed considering the population as the element
203 at risk. To account for population growth, a multi-temporal risk assessment has been executed
204 considering data representative of the population distribution over the study area in 2010, 2018, 2019
205 and 2020 (the only available years). Although population data and their future projection are available
206 in absolute terms for the city of Cuenca, the spatial distribution over the municipalities forming the
207 city of Cuenca is not available. Such datum is notoriously essential in any landslide risk evaluation.
208 For this work, distribution of energy supply contracts data, derived from the IRSE (*Instituto de*
209 *Estudios de Régimen Seccional del Ecuador* - Institute of Studies on the Sectional Regime of Ecuador,

210 <http://ierse.uazuay.edu.ec/>), were used. Although it does not correspond to the number of people
 211 living in each sector of the city, the provision of utilities (i.e., electricity supply contracts) is a proxy
 212 of the population and can be considered as alternative data in relative risk assessment perspective.
 213 Since energy supply contracts data consists of vector points, for each municipality such information
 214 were aggregated and associated to the specific area and used for relative risk evaluation. Figures 4a,
 215 b, c and d depict the distribution of such data at the municipalities scale between 2010 and 2020 and
 216 their variation in comparison with the year 2010 (Fig. 4 e, f, g).



217
 218 Fig. 4. a), b), c) and d) Multi-temporal municipalities evolution by analysing electricity supply contracts from 2010 to
 219 2020; e), f) and g) Variation of power supply contracts, respectively in years 2018, 2019 and 2020, compared to 2010.

220

221 3.3. Susceptibility analysis

222 Based on the mapped distribution of landslides and as a first modeling step in the evaluation of the
 223 relative landslide risk in the city of Cuenca, landslide susceptibility was estimated using a machine

224 learning algorithm (Lombardo et al. 2016; Di Napoli et al. 2020b, 2021). In this perspective,
225 environmental covariates acting as potential predisposing factors for landslide initiation were selected
226 and tested for multicollinearity. Subsequently, covariates that do not exhibit multicollinearity were
227 used for susceptibility estimation through the MaxEnt algorithm (Phillips et al. 2006; Phillips and
228 Dudík 2008). Results were validated using multiple criteria. Details about evaluation steps are
229 provided below.

230

231 **3.3.1. Covariates selection and multicollinearity analysis**

232 For the analysis eleven covariates such as *i*) Slope Steepness; *ii*) Eastness; *iii*) Northness; *iv*) Planar
233 and *v*) Profile Curvatures; *vi*) Topographic Wetness Index (TWI); *vii*) Relative Slope Position (RSP);
234 *viii*) Distance to streams; *ix*) Distance to roads; *x*) Land Use; and *xi*) Geology were selected
235 (Supplementary Material 1). Numerical covariates were derived from a 10 × 10 m Digital Elevation
236 Model resampled from an original 3 × 3 m Digital Terrain Model. DTM resampling and topography-
237 related covariates raster generation were completed into a GIS environment (i.e., SAGA GIS, Conrad
238 et al., 2015). Categorical covariates such as Land use and Geology were derived considering data
239 available from the National Institute of Geology and Energy (<https://sni.gob.ec>).

240 In general, the likelihood of a landslide occurrence is positively correlated with slope due to its effect
241 in modulating acting force and slope aspect. The aspect is radial in nature, with values 360 and 1
242 being adjacent degree measurements. A common way to treat radial data is to transform them by
243 using trigonometric functions. A trigonometric transformation of aspect data is rather “pure” since it
244 retains the continuity of aspect. For these data, a *cos* transformation measures southerliness-to-
245 northerliness (-1 to 1, respectively), while a *sin* transformation measures westerliness-to-easterliness
246 (-1 to 1, respectively), obtaining northness and eastness (Lombardo et al. 2020). Planform curvature
247 relates to the convergence and divergence of flow across a surface, so that it is a proxy of potential
248 runoff concentration. Profile curvature affects acceleration and deceleration of runoff across the
249 surface indicating the predisposition of a slope to soil erosion. Topographic Wetness Index is an

250 important factor indicating the potential of runoff generation and is a proxy for the thickness of the
 251 saturation zone. Shallow landslides are facilitated by soil saturation. High index values indicate the
 252 great potential of water accumulated due to low slope angles. Relative slope position indicates the
 253 location of each cell relative to the ridge and valley of a hillslope. Distance to streams and distance
 254 to roads have been both estimated by using the Euclidean distance method. Distance to stream is a
 255 crucial parameter that controls slope stability. In fact, slope foot erosion due to stream water flow is
 256 a common triggering factor for landslide initiation. Similarly, landslide initiation can be facilitated
 257 by the presence of roads due to runoff concentration and preferential infiltration and the presence of
 258 weak material due to excavation. Land use provides information about the potential practice that
 259 might favour landslide development. Geology of a slope is a significant predisposing factor for
 260 landslide initiation since properties of slope materials effectively control initiation potential. For
 261 categorical covariates, raster cells codes were assigned according to Table 1.

262
 263 Table 1. Land use and geology types' code description.

<i>Variable</i>	<i>Code</i>	<i>Class</i>	<i>Variable</i>	<i>Code</i>	<i>Class</i>	<i>Code</i>	<i>Class</i>
<i>Land Use</i>	<i>1</i>	Bare area	<i>Geology</i>	<i>1</i>	PT – Tarqui Fm.	<i>10</i>	Kc – Celica Fm.
	<i>2</i>	Forest		<i>2</i>	gt – Alluvial deposit	<i>11</i>	MAz – Azogues Fm.
	<i>3</i>	Crop		<i>3</i>	PIR – Santa Rosa Fm.	<i>12</i>	Tv – Travertine
	<i>4</i>	Moor		<i>4</i>	PTu – Turi Fm.	<i>13</i>	T – Fluvial terracce
	<i>5</i>	Urban area		<i>5</i>	MM – Mangan Fm.	<i>14</i>	QLI – Llacao Fm.
	<i>6</i>	Shrub vegetation		<i>6</i>	MB – Biblian Fm.	<i>15</i>	Dc – Colluvial deposit
	<i>7</i>	Water course		<i>7</i>	Da – Alluvial deposit	<i>16</i>	K7 – Yunguilla Fm.
	<i>8</i>	Corn cultivation		<i>8</i>	ML – Loyola Fm.	<i>17</i>	Av – Varvada clays
	<i>9</i>	Natural grassland		<i>9</i>	De – Colluvial deposit		

264
 265 Once selected, covariates were tested for multicollinearity. Multicollinearity represents the
 266 occurrence of high intercorrelations among two or more independent variables within a predictive

267 model and can lead to skewed or misleading results. To identify multicollinearity among selected
268 variables the Variance Inflation Factor (VIF) through the “usdm” package (Naimi et al. 2014) was
269 employed. VIF measures how much of the variation in one variable is explained by the other variable.
270 It estimates how much the variance of a coefficient is "inflated", for this reason VIF, because of linear
271 dependence with other predictors. VIF can be calculated by using the formula $1/(1-R^2)$, where R^2 is
272 the coefficient of determination of the regression equation. The smallest possible value of VIF is one
273 (absence of multicollinearity). As a rule of thumb, a VIF value that exceeds 5 or 10 indicates a
274 problematic amount of collinearity (Gareth et al. 2013). However, VIF values greater than 2.50 should
275 be treated with caution since they correspond to an R^2 of 0.60 with the other variables. When faced
276 with multicollinearity, the concerned variables should be removed, since the presence of
277 multicollinearity implies that the information provided about the response by this variable is
278 redundant when other variables are present (Gareth et al. 2013).

279

280 **3.3.2. Susceptibility assessment**

281 Landslide susceptibility was assessed using the MaxEnt modelling algorithm and identified
282 covariates (Elith et al. 2011). MaxEnt is a presence-only (PO) spatial distribution method that deals
283 only with landslide presence locations (Zhao et al. 2020). It makes use of occurrence data and a large
284 (typically 10000) number of points throughout the study area, which are referred to as background
285 points. Background points define the frequency distribution of available environmental variables in
286 the landscape. To reconstruct the potential distribution of an event, MaxEnt calculates two probability
287 densities. For all presence points, probability density describes the relative likelihood of all
288 environmental variables in the model over the range of those points, describing the environment
289 where an event has occurred. Then, the algorithm calculates the density of landslide occurrences
290 across the entire landscape, based on the background points that characterize the available
291 environment within the study region. Population size is typically unknown, so only relative
292 comparisons among these rates are meaningful, resulting in a Relative Occurrence Rate (ROR;

293 Fithian and Hastie, 2013). ROR can be seen as the ratio between the probability density of covariates
294 across locations within the considered geographic space where the landslide is present and the
295 probability density of covariates across the entire geographic space, thus obtaining insights on the
296 relative proneness to failure of a given cell compared to another one: the map of probability of event
297 occurrence ranges from 0 (i.e., no landslide probability) to 1 (highest landslide probability).
298 Since MaxEnt predictions are sensitive to initial modelling settings (Merow et al. 2013), different
299 MaxEnt implementations were evaluated through the ENMeval R package (R Core Team 2021) to
300 detect the settings that optimize the trade-off between goodness-of-fit and overfitting (Muscarella et
301 al. 2014). In fact, MaxEnt is possible to set up two main parameters: 1) feature classes and 2)
302 regularization multiplier. Feature class represents a mathematical transformation of the different
303 covariates used in the model to allow complex relationships to be modelled (Elith et al. 2010). The
304 regularization multiplier is a parameter that adds new constraints, in other words, is a penalty imposed
305 on the model. The main goal is to prevent over-complexity and/or overfitting by controlling the
306 intensity of the chosen feature classes used to build the model (Elith et al. 2010). For a detailed
307 explanation of feature classes and regularization multipliers, it is recommended to consult Merow et
308 al., (2013). For the analysis, regularization values between 0.5 and 10, with 0.5 steps were
309 investigated, and the following feature classes were considered: linear, linear + quadratic, hinge,
310 linear + quadratic + hinge, linear + quadratic + hinge + product, and linear + quadratic + hinge +
311 product + threshold (Muscarella et al. 2014).

312

313 **3.3.3. Model evaluation and predictor contribution**

314 Model evaluation was completed using spatial block cross-validation scheme (Muscarella et al. 2014)
315 implemented in ENMeval. This method converts part of occurrence records and background points
316 into evaluation bins and uses them to reduce spatial - autocorrelation between training and validation
317 points that can overinflate model performance in presence of biased sampling (Hijmans 2012; Wenger
318 and Olden 2012). The block cross-validation scheme proved able to assess model transferability, i.e.,

319 the ability to extrapolate predictions into new areas (Roberts et al. 2017) and to penalize models based
320 on meaningless predictors (Fourcade et al. 2018).

321 Because no consensus currently exists regarding the most appropriate metric or approach to evaluate
322 the performance of models (Fielding and Bell 1997; Warren and Seifert 2011; Peterson et al. 2011),
323 different statistical approaches have been adopted to assess the models' predictive performance with
324 presence-background data (Muscarella et al. 2014). In this case, the best model reliability-
325 combination has been chosen following three criteria: *i*) lowest delta Akaike Information Criteria
326 (ΔAIC_c) (Burnham and Anderson 2002), *ii*) Area Under the Curve plot based on the training data
327 (AUC_{train}) (Hanley and McNeil 1982) and *iii*) the difference between training and testing AUC
328 (AUC_{diff}) (Warren and Seifert 2011).

329 AIC_c is calculated using the full data set and its metrics are not affected by the method chosen for
330 data partitioning. AIC is a single number score that can be used to determine which of multiple models
331 is most likely to be the best model for a given dataset. It estimates models likelihood in a relative
332 manner, meaning that AIC scores are only useful in comparison with other AIC scores for the same
333 dataset. A lower AIC score indicates higher model performance. AIC is most frequently used in
334 situations where one is not able to easily test the model's performance on a test set. Furthermore, AIC
335 results are reported as ΔAIC_c scores because it is the easiest way to calculate and interpret them. The
336 ΔAIC_c is the relative difference between the best model and each other model in the dataset. The
337 formula is the following (Eq. 1):

$$338 \quad \Delta AIC_c = AIC_i - \min AIC \quad (1)$$

339 where:

340 AIC_i is the score for the particular model i , and $\min AIC$ is the score for the "best" model.

341 Hence, AIC values closely to zero or equal to zero indicate the best model with the available dataset.

342 The AUC is the measure of the ability of a classifier to distinguish between classes and is used as a
343 summary of the Receiver Operator Characteristic (ROC) curve, which is an evaluation metric for
344 binary classification problems. A high AUC, which ranges between 0 and 1, indicates that sites with

345 high predicted suitability values tend to be areas of known presence, while locations with lower model
346 prediction values tend to be areas where the landslide is not known to be present (absent or random
347 point). Lastly, to quantify overfitting, ENMeval calculates the difference between training and testing
348 AUC (AUC_{diff}), which is expected to be high with overfit models.

349 Moreover, the Landslide Ratio of each predicted landslide susceptibility class (LR_{class}) has been
350 employed as a further performance evaluation of the landslide model. LR_{class} is based on the ratio of
351 the number of unstable sites contained in each susceptibility class, in relation to the total number of
352 actual landslide sites, according to the predicted percentage of area in each class of susceptibility
353 category (Eq. 2). This index was developed specifically to deal with situations when boundaries of
354 observed landslides are not available, but where their locations are known. The advantage of using
355 LR_{class} index is that it considers both the predicted stable and unstable areas and thus significantly
356 decreases over-prediction.

$$357 \quad LR_{class} = \frac{\% \text{ of contained sites in each susceptibility class}}{\% \text{ of predicted landslide areas in each susceptibility class}} \quad (2)$$

358 LR_{class} index indicates that if a slope failure occurs, the predicted unstable area has a chance equal to
359 LR_{class} of including an actual slope failure. A larger value of LR_{class} corresponds to a lower over-
360 prediction by the model (Park et al. 2013).

361 Predictive performance estimation is only a partial metric of model goodness. Predictor contributions
362 represent a further key step that should be assessed to comprehensively estimate the validity of a
363 model for relating results to the analyzed processes. In this contribution, the investigation has been
364 carried out considering 1) predictor importance and 2) percentage contribution (Oke and Thompson
365 2015). Predictor importance represents the degree to which single environmental variables are
366 contributing to the final model, so that the percent contributions for all predictors in a model sum to
367 100% (Phillips 2008). The percentage contribution, called permutation importance, is determined by
368 randomly altering the values of that variable among the training points (both presence and
369 background) and measuring the resulting decrease in training AUC. A large decrease indicates that

370 the model depends heavily on that variable (Phillips 2017). A very useful and detailed explanation
371 was given by Bradie and Leung (2017).

372

373 **3.4. Relative risk analysis**

374 Considering the definition of risk introduced by Varnes (1984) as “the expected number of lives lost,
375 persons injured, damage to property and disruption of economic activity due to a particular damaging
376 phenomenon for a given area and reference period”, landslide risk can be assessed qualitatively
377 (Wang et al. 2013) or quantitatively (Chang et al. 2021). Generally, for a wide area, where the quality
378 and quantity of available data are inadequate for quantitative analysis, a qualitative risk evaluation
379 may be more appropriate (Andrejev et al. 2017). In the context of the proposed analysis, data from
380 susceptibility analysis and electricity supply contracts were used as a basis for multi-temporal
381 landslide risk evaluation over the study area. The evaluation was completed considering only the risk
382 for people. As for landslide susceptibility, results of the estimation derived by the described Machine
383 Learning-based approach were employed. As for people at risk quantification, the number of power
384 supply contracts and their relative variation between 2010 and 2020 (i.e. 2010, 2018, 2019 and 2020)
385 were considered in the assumption that the number of contracts is a good proxy of people distribution
386 over the study area. The use of this proxy well fits the choice of evaluating the risk due to landslide
387 in a relative manner. Indeed, relative risk is considered as the intersection of the landslide
388 susceptibility and the number of elements exposed to landslides (Andrejev et al. 2017). In the case of
389 the city of Cuenca, the number of elements at risk (i.e. people) is not fully known in terms of spatial
390 distribution and the presence of electricity supply contracts location, indicating the number of groups
391 of people for each location, represents somehow an opportunity to overcome this issue. Since both
392 landslide susceptibility and power supply contract are georeferenced, but supply contracts are not
393 regularly distributed, the relative risk assessment was completed at the scale of the municipalities of
394 the city. Obtained outcomes in the form of relative risk histograms, depicting the number of electricity
395 supply contracts located in each susceptibility class, were classified using the Sturges method

396 (Sturges 1926), which allows highlighting how over time the different areas of the city of Cuenca
397 underwent an increase in contracts and therefore also in relative risk.

398

399 **4. Results and discussion**

400 **4.1 Landslide inventory map**

401 The obtained LIM is composed of 710 landslides detected through different approaches and all
402 validated thanks to field surveys. Landslide database contains useful information regarding the type
403 of movement according to Cruden and Varnes (1996), state of activity, location, triggering factor (i.e.,
404 precipitation or anthropic), geology, land use, velocity and further information. In the database, it is
405 possible to recognize rockfalls (72 – 10.1%), topples (3 – 0.4%), flows (8 – 1.1%), spreads (5 – 0.7%),
406 rotational slides (550 – 77.1%) and translational slides (72 – 10.1%) (Fig. 5a). These phenomena
407 represent the principal hazard of the area, since they affect the urban area damaging roads networks
408 and buildings. The main causes of landslides triggering are intense or prolonged rainfalls and mining
409 activity (i.e. incorrect management of the excavation face; Jaboyedoff et al., 2016).

410 Geostructural aspects notably influence the occurrence of gravitational phenomena involving rock
411 masses such as falls, topples and planar slides. Rockfalls and topples affect steep artificial slopes
412 around the main infrastructures (Miele et al. 2021). In fact, these phenomena mainly occur where
413 anthropic actions have provoked cutting linked to the construction of infrastructures. The high slope
414 angles represent an essential element in favouring translational slides (Raso et al. 2020), whose action
415 is often enhanced by diffuse and channelled erosion operated by running water.

416

417 **4.2 Multicollinearity examination**

418 Table 2 shows the results of multicollinearity analysis carried out through VIF estimation and its
419 comparison with a predefined threshold value. In this regard, it must be noted that there is no
420 unequivocal and approved threshold in the scientific literature. However, it is generally accepted that

421 VIF values higher than 10 indicate severe collinearity (Hair et al. 2010), even though this rule of
 422 thumb lacks a theoretical basis (Gómez et al. 2016).

423 The employment of many environmental covariates might lead to overfitting problems, but, in this
 424 work, the individual predisposing factor values differ considerably from the aforementioned
 425 threshold. Based on Table 2, the highest VIF value is 2.02, corresponding to the Topographic Wetness
 426 Index, while the smallest ones are 1.01 and 1.02, which are associated with Northness and Eastness,
 427 respectively. Accordingly, there are no environmental variables that exceed the critical value, and
 428 thus, these results satisfy the criterion ($VIF < 5$) proving that there is no multicollinearity among the
 429 landslide PFs.

430 Table 2. Multicollinearity analysis for the landslide environmental factors.

<i>Environmental Variables</i>	<i>Variance Inflation Factor</i>
<i>Slope steepness</i>	1.72
<i>Eastness</i>	1.02
<i>Northness</i>	1.01
<i>Planar curvature</i>	1.39
<i>Profile curvature</i>	1.30
<i>Topographic Wetness Index</i>	2.02
<i>Relative Slope Position</i>	1.39
<i>Distance to stream</i>	1.12
<i>Distance to road</i>	1.19

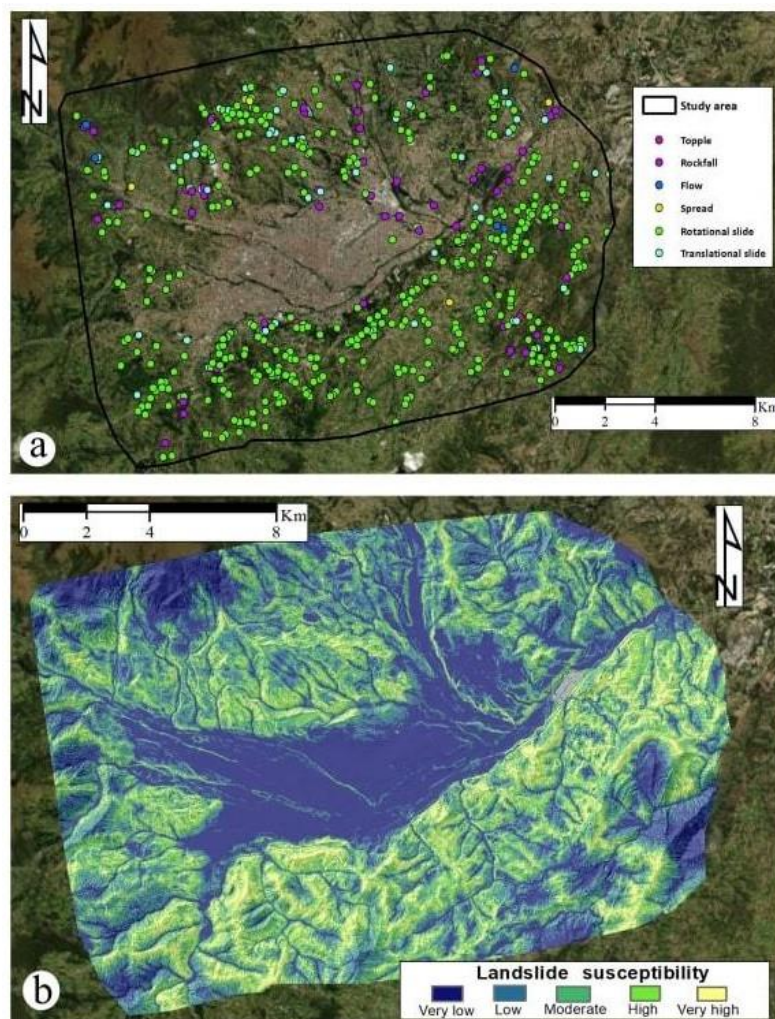
431

432 **4.3 Landslide susceptibility**

433 Figure 5b shows the result of landslide susceptibility analysis in the form of a susceptibility map
 434 subdivided into five classes through Natural Breaks distribution (Jenks 1967). Each class represents
 435 a specific susceptibility range including very low, low, moderate, high and very high susceptibility
 436 levels. Natural Breaks classification, also called Jenks optimization method, is a data classification
 437 method designed to determine the best arrangement of values into different classes. This is done by
 438 seeking to reduce the standard deviation value within each class and maximizing that between the

439 classes themselves (Basofi et al. 2018; Novellino et al. 2021). The percentage of susceptibility classes
440 is summarized in Table 3.

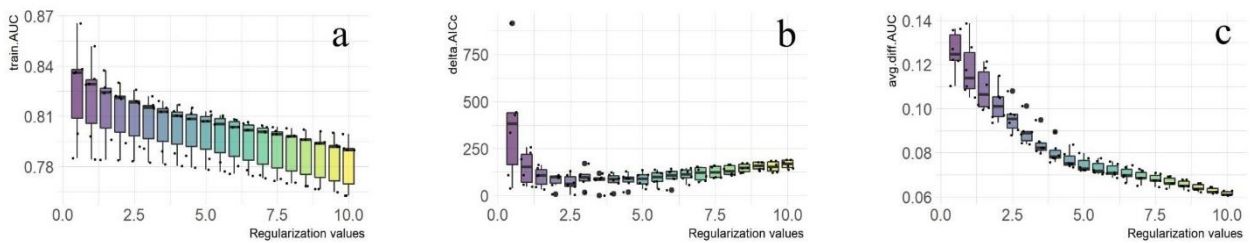
441 As reported in the map, the most susceptible areas of the city of Cuenca are observable at
442 mountainsides that border the city. In these areas, slopes are steep and concave, and roads create local
443 discontinuities. The central part of the map is characterized by very low and low susceptibility zones,
444 and represent the Cuenca urban area, located in the plain which is characterized by the presence of
445 alluvial deposits and different terraces orders. Along the *Tomebamba* shores, due to the erosive action
446 of the rivers affecting the foot of the slopes, there are areas predisposed to landsliding with a medium
447 to very high susceptibility. Periurban and rural areas, instead, are in medium to very high
448 susceptibility areas where there are steeper mountainsides.



449
450 Fig. 5. a) Landslide inventory map of Cuenca. Landslides are represented as points (data source modified from Miele et
451 al. (2021); b) landslide susceptibility map of the study area obtained by means of MaxEnt algorithm implementation.

452 **4.4 Susceptibility model validation**

453 According to Swets (1988), the obtained models have achieved fair-to-good predictive performance,
454 with AUC values ranging from 0.763 to 0.866 (Fig. 6). The lower value is associated with models
455 with linear or linear + quadratic features and high regularization values (i.e., 9.5 and 10). The higher
456 value is associated with a model with all features and low regularization values (i.e., 0.5 and 1).
457 AUC_{diff} values scored from 0.06 to 0.14. Among the resulting 120 combinations, the one reporting
458 the lowest ΔAIC_c has been chosen. The selected model is characterized by the following peculiarity:
459 linear + quadratic + hinge + product + threshold features, AUC value of 0.82, average AUC difference
460 value of 0.08 and ΔAIC_c value equal to 0 (for further information refer to Supplementary Material 2).



461

462 Fig. 6. a) Boxplot of AUC training data ($AUC_{train} - \text{train.AUC}$); b) the lowest difference between the best model and
463 each other model in the dataset ($\Delta AIC_c - \text{delta.AIC}_c$); c) difference between training and testing AUC ($AUC_{diff} -$
464 avg.diff.AUC).

465 The availability of a LIM has made possible the evaluation of model performance also considering
466 field data. Intersecting the landslide detachment points and the final susceptibility map, it has been
467 possible to achieve information about landslide distribution and areal extent of the susceptibility
468 classes. Areas characterized by high and very high susceptibility involve 15.5% and 10.6% of the
469 total study area, respectively (Table 3). Very low and low susceptibility classes cover about 55% of
470 the study area, falling into the central sectors, namely Cuenca city. The remaining portions are
471 assigned to the moderate (19.4%) class. Moreover, the highest concentration of landslides can be
472 found within the highest susceptibility class values (high, 25.2%, and very high, 48.9%). In addition,
473 about 3% falling into the very low susceptibility class and the remaining 22.9% are distributed in the
474 low class (i.e., 9.4%) and moderate class (i.e., 13.5%). The last column of Table 3 highlights the

475 LR_{class} percentage for each susceptibility class. As it can be noted, the highest LR_{class} value
 476 corresponds to the very high susceptibility class and more than 80% falling into high and very high
 477 susceptibility values. This evidence roughly implies that, if a landslide occurs, then the predicted
 478 susceptible area has about an 80% chance of including the landslide itself.

479

480 Table 3. Summary of Maxent outcomes in landslide simulations.

<i>Susceptibility classes</i>	<i>Landslide site (a)</i>	<i>% of landslide site (c)=a/b</i>	<i>% of predicted area (d)</i>	<i>LR_{class} (e) = c/d</i>	<i>% of LR_{class} = e/f</i>
<i>Very low</i>	14	3.0	30.6	0.1	1.4
<i>Low</i>	44	9.4	23.9	0.9	5.3
<i>Moderate</i>	63	13.5	19.4	0.7	9.4
<i>High</i>	118	25.2	15.5	1.6	21.9
<i>Very high</i>	229	48.9	10.6	4.6	62.0
<i>Sum</i>	468 (b)	100	100	7.9 (f)	100

481

482 The spatial aggregation of the susceptibility map confirmed that the largest part of the study region
 483 has a low susceptibility to the occurrence of landslide events. Therefore, results highlight that almost
 484 75% of actual landslides were localized in the high and very high susceptibility classes. Also, the
 485 higher susceptibility classes showed higher values of LR_{class} percentages. These outcomes show
 486 significant agreement in quantitative terms between the simulated scenario and landslides inventory
 487 map. All the produced analysis permits zoning the complex territory of the Cuenca area to identify
 488 the spatial probability of landslides initiation in areas characterized by specific conditions
 489 materialized by the considered environmental variables. Moreover, landslide distribution is
 490 characterized by an increasing trend when passing from the lowest to the highest classes of
 491 susceptibility. These observations highlight that, despite the limited area extension of the very high
 492 susceptibility class, most of the landslides surveyed fall within the latter. Furthermore, the results of
 493 this elaboration have made clear the need for preventive action, perhaps based on simple monitoring
 494 techniques, to avoid the worsening of local geoenvironmental conditions.

495

496 **4.5 Factors predisposing slope instability**

497 As the latest outcome, the variables contribution has been accomplished. This result allows
 498 understanding which variables have greater importance in the final models' implementation. Table 4
 499 presents the impact of each variable. In particular, the results reveal that the highest conditioning
 500 variables (i.e. the variables that assume a fundamental role in the final landslide susceptibility map)
 501 are slope steepness, distance to roads and planform curvature. So this means that a model with a
 502 higher fit is achieved through the aforementioned variables. In general, when percent contribution
 503 was observed it is desirable to see a nice spread of values. Conversely, if the contribution of a variable
 504 is high in the model (i.e. higher than 70%) something is not right and that variable is not encompassing
 505 many variations or that variable is correlated with a bunch of other variables. Other variables, such
 506 as distance to streams, land use, geology and relative slope position show noteworthy values in the
 507 model. Lastly, low values close to zero are assigned to eastness, northness, profile curvature and
 508 Topographic Wetness Index.

509 Table 4. Variable contribution values of environmental factors.

<i>Environmental Variables</i>	<i>Percent contribution</i>	<i>Permutation importance</i>
<i>Slope steepness</i>	26.6	26.6
<i>Eastness</i>	1.7	0.7
<i>Northness</i>	0.1	0.4
<i>Planar curvature</i>	19.5	15.7
<i>Profile curvature</i>	0.8	0.8
<i>Topographic Wetness Index</i>	0.2	0.1
<i>Relative Slope Position</i>	6.1	1.5
<i>Distance to stream</i>	6.0	5.5
<i>Distance to road</i>	20.3	26.6
<i>Geology</i>	9.1	8.2
<i>Land use</i>	9.4	10.6

510

511 Primary roles in the slope stability are related to slope steepness, which always influence the water
512 infiltration, upslope flow intensity and gravity force effect on safety factor against slope instability
513 (Huat et al. 2006). A particular condition is related to the presence of road cuts that influences water
514 infiltration and flows as well due to impervious pavement surface. An additional fundamental
515 covariate contribution on this zone's stability is represented by the planar curvature that adjusting the
516 convergence or divergence of water in the direction of landslide movement and landslide material
517 (Ohlmacher 2007). Furthermore, in the Cuenca territory, hillsides with planar curvature are the most
518 susceptible to earth and debris flows, and earth and debris slides. Indeed, this factor is used to identify
519 gullies (Wieczorek et al. 1997), and debris flow initiation areas can be recognized where the curvature
520 values are negative (Park et al. 2016). Finally, another important predisposing action to slope stability
521 is related to geology. Unconsolidated material such as alluvial, colluvial deposits or pyroclastic
522 lithologies deriving from the near volcanoes' activity cover the surrounding mountainous landscape,
523 determinating a very high susceptibility to sliding.

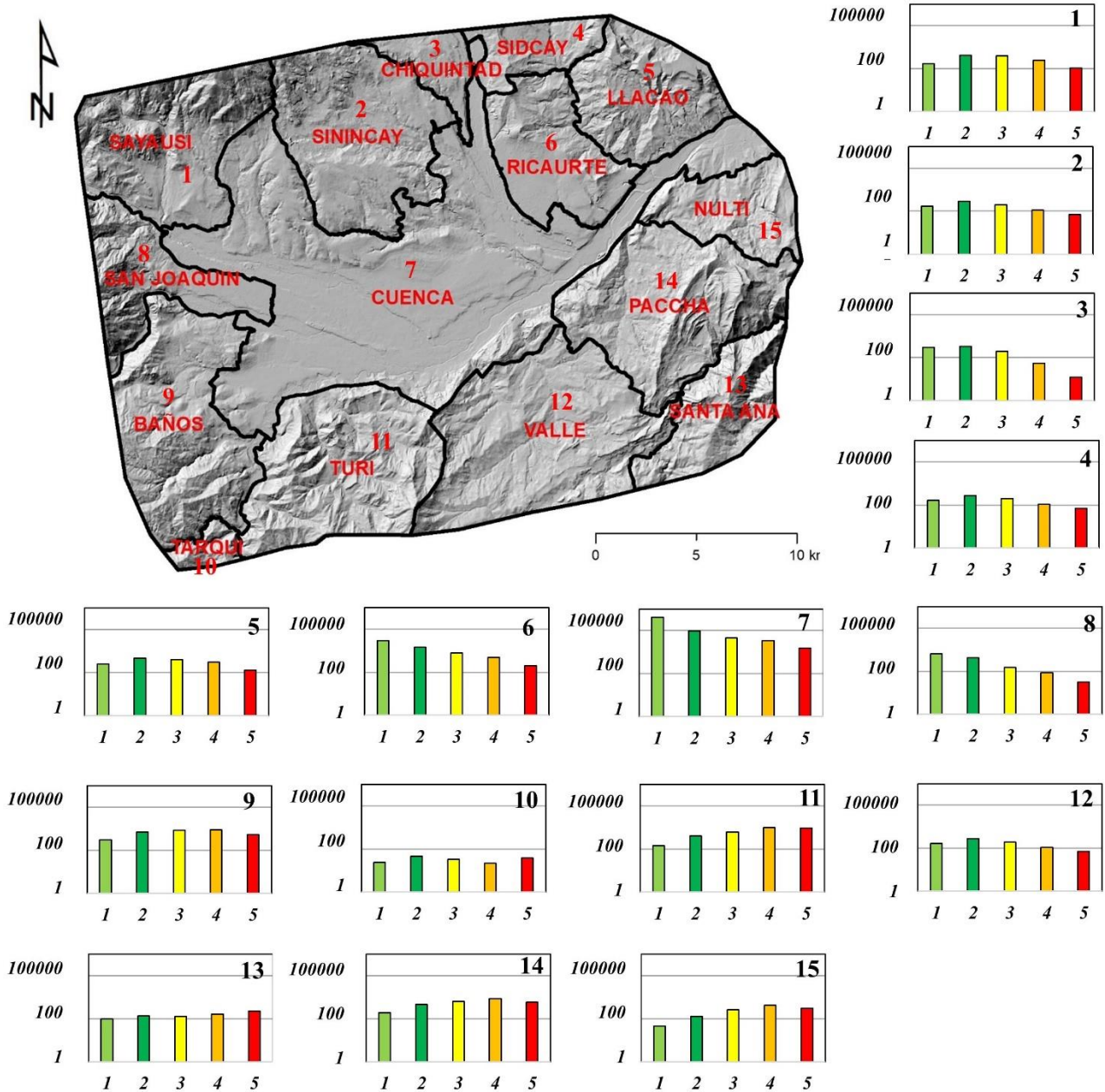
524

525 **4.6 Relative risk assessment**

526 Figure 7 provides an overview of relative risk in each municipality updated to 2020. As observable
527 from the inset graphs, the relative risk is higher in boroughs that surround the southern portions of
528 Cuenca, namely *Turi*, *Valle*, *Santa Ana*, *Tarqui* and *Paccha*. Such boroughs show a higher number
529 of energy supply contracts (i.e. element at risk) in high and very high landslide susceptibility classes.
530 On the contrary, the northern sectors of the study area report a lower number of power supply
531 contracts included in the higher landslide susceptibility class. Lastly, the central portion of the study
532 area, in which the city of Cuenca falls, is characterized chiefly by very low and low landslide risk
533 values. However, recently, the city of Cuenca has experienced a substantial demographic increase
534 which has led to the construction of buildings in notoriously very high susceptibility areas. Similar
535 problems can be easily found in other rapidly expanding cities where the construction of new
536 boroughs takes place in the hilly and mountainous areas that are more prone to instability (Di Martire

537 et al. 2012). This condition is evidenced by the high number of electricity supply contracts falling
 538 into the highest susceptibility classes.

539



540

541 Fig. 7. Landslide risk outcome. The graphs, on a logarithmic scale, highlights the electricity supply contracts numbers
 542 for each risk class (1: very low; 2: low; 3: moderate; 4: high; 5: very high). The numbers in the upper-right corner of the
 543 graphs represent the municipalities of the studied area.

544 Since the city has experienced consistent growth of population between 2010 and 2020 with a
 545 sustained occupation of peri-urban hilly areas characterized by higher susceptibility to landslide in
 546 comparison to the center of the city, a specific exploration of the spatial and temporal rate of change

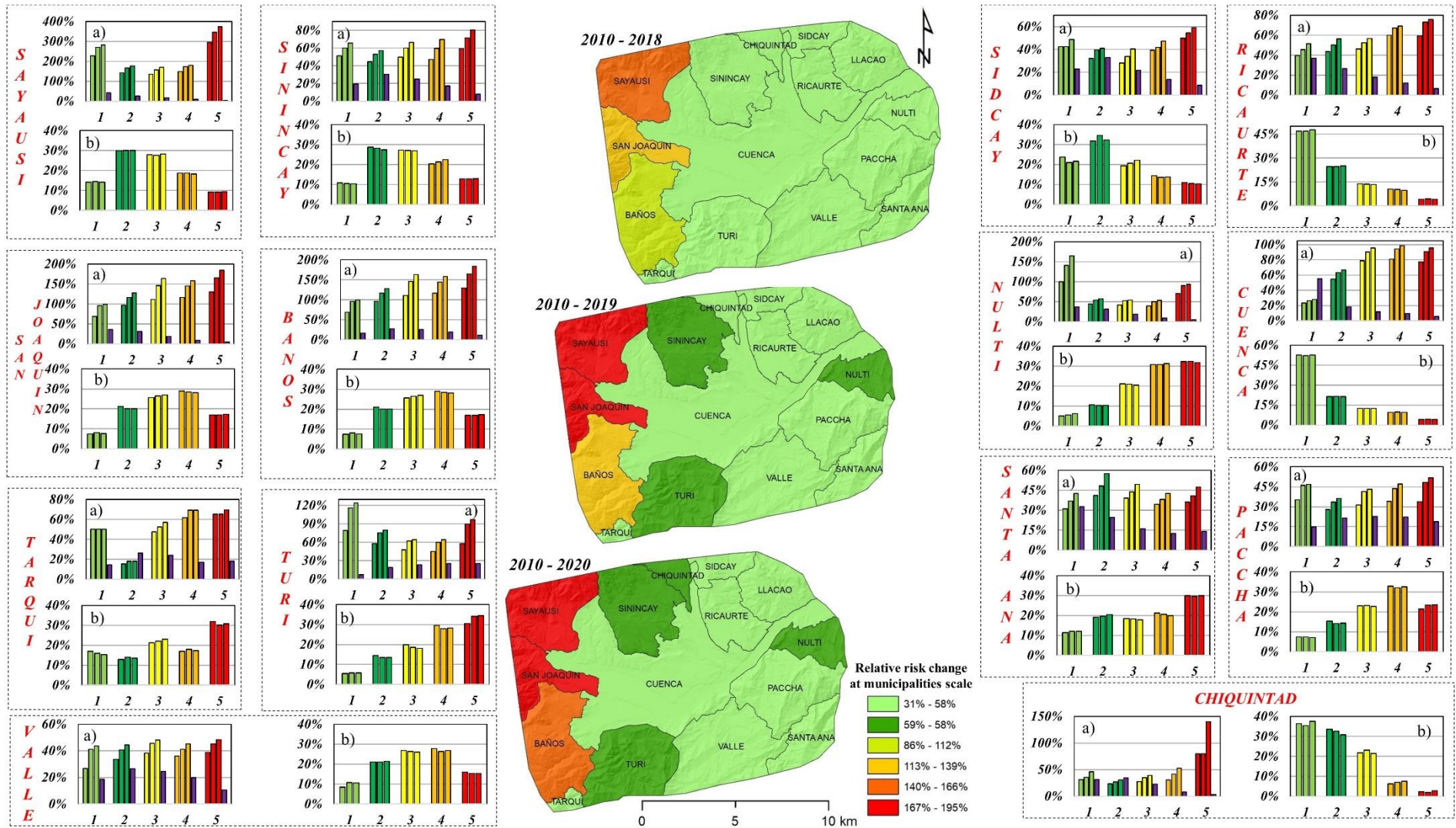
547 of both electricity supply contracts and relative risk for each municipality is reported in Figure 8. In
548 particular, in Figure 8, graph “a” highlight electricity supply contracts for each landslide susceptibility
549 class obtaining percentage of change of the relative landslide risk from 2010 to 2020, while the same
550 rate in terms of total risk change (i.e. normalized to the total 100%) is reported in graph “b”. In
551 addition, the relative areal extension of different susceptibility classes is also reported on graph “a”
552 (violet bar). In the reference period, the municipality of *Sayausi*, *San Joaquin*, *Tarqui*, *Valle*, *Sidcay*,
553 *Banos*, *Ricaurte*, *Paccha* and *Chiquintad* experienced a higher increase in electricity supply contacts
554 located in very high susceptibility class corresponding to an increase in relative landslide risk (Fig.
555 8). Such an increase, in comparison with 2010, ranged between 33% of *Paccha* and 300% of *Sayausi*.
556 Conversely, the municipalities of *Turi*, *Nulti*, *Santa Ana* and Cuenca experienced a higher increase in
557 electricity supply contacts located in very low to medium susceptibility class corresponding, in most
558 of the cases, to a decrease in relative landslide risk. Such an increase, in comparison with 2010, ranged
559 between 20% of Cuenca and 100% of *Nulti*.

560 However, in absolute terms, the districts of *Tarqui*, *Turi*, *Nulti* and *Santa Ana* show the highest
561 relative risk with the highest number of electricity supply contracts located in high susceptibility areas
562 between 2018 and 2020. Conversely, the boroughs of *Sayausi*, *Sidcay*, *Ricaurte*, Cuenca and
563 *Ciquintad* show the lowest relative risk with the lowest number of electricity supply contracts located
564 in high susceptibility areas in the same period. In the supplementary materials (Supplementary
565 Material 3), it is possible to consult the tables that quantitatively represent the graphs shown in Figure
566 8.

567 Landslide susceptibility is a widely used tool to assess the areas most prone to instability. In the last
568 decade, these analyzes have also been conducted in emerging and developing countries (O’Hare and
569 Rivas 2005; Klimeš and Rios Escobar 2010; Listo and Carvalho Vieira 2012; Jamalullail et al. 2021).
570 Several attempts have been made to estimate landslide risk in various contexts where demographic
571 growth is very pronounced (Rahman 2012; Listo and Carvalho Vieira 2012; Rojas et al. 2013;
572 Alcántara-Ayala and Moreno 2016). Considering the high growth rate recorded in the last few years,

573 it is appropriate to assess the exposure of the landslide risk over time. This type of analysis made it
574 possible to identify the most critical areas and sectors of the city of Cuenca also in terms of risk
575 evolution due to population growth. The outcome of the analysis represents a significant land
576 planning tool for the definition of urban occupation plans, land-use prescriptions, and mitigation
577 protocols that should be applied to reduce the impact of a landslide occurring in urban areas (Klimeš
578 et al. 2020; Sultana and Tan 2021). The problem of landslides involving settlements and claiming
579 human lives is of particular significance in low to mid-income countries because people are often
580 concentrated in high-hazard urban areas and vulnerability factors like poorly building of housing,
581 poor governance and the lack of experience in facing natural disasters and the absence of early
582 warning systems consistently exacerbate landslide impact (Petley 2012; Zorn 2018; Aguirre-Ayerbe
583 et al. 2020).

584 This qualitative procedure for evaluating the landslide exposure in Cuenca tries to provide
585 information for risk assessment, useful in a preliminary stage of regional planning or for more detailed
586 studies on the high-exposure areas. Therefore, the procedure proposed in this study could be
587 implemented when not all the information useful for the risk assessment are attainable. Exposure
588 quantification, which is a basic input in spatial and risk reduction planning, is the main objective of
589 this study. It is important to mention that, as not all the information are available, landslide risk values
590 are not expressed in absolute terms, but relative landslide risk could be a good proxy of districts that
591 have encountered, in the last decade, a population growing falling into very high-risk territories.
592 Notwithstanding the limitations, this study has allowed estimating which areas are more prone to
593 instability, which areas have a high relative landslide risk and also to establish the changes of risk in
594 future by consulting the trend in the different municipalities.



595

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597

598

Fig. 8. Multi-temporal evolutionary perspective of relative risk to a landslide. Graphs “a” highlight electricity supply contracts for each landslide susceptibility class; graphs “b” represent the same rate in terms of total risk, namely normalized to the total 100%. The coloured bars of the histograms indicate the different risk classes ranging from very low to very high, the violet bar, included in the histogram “a”, specifies the landslide susceptibility classes areal extension.

599 **5. Conclusions**

600 In this paper, an analysis of relative landslide risk, and its multi-temporal variation between 2010 and
601 2020, for the city of Cuenca in Ecuador (Latin America) has been presented. This study provided
602 insights into important issues such as *i*) the effect of the sustained expansion of urban areas due to
603 population growth on relative landslide risk variation and *ii*) reduced complexity method of risk
604 assessment in the presence of partial data only (i.e. landslide susceptibility rather than hazard, and
605 electricity supply contracts rather than population distribution). Results indicate that current higher
606 relative risk is estimated for districts located at the southern sector of the study area (i.e. *Turi, Valle,*
607 *Santa Ana, Tarqui* and *Paccha*). In addition, the multi-temporal analysis indicates that most boroughs
608 of the city located in the hilly areas that bound the center (i.e. *Sayausi, San Joaquin, Tarqui, Valle,*
609 *Sidcay, Banos, Sidcay, Ricaurte, Paccha* and *Chiquintad*), experiencing sustained population growth,
610 will be exposed to an increased risk with a sustained growth trend. This is also connected to the
611 overall high vulnerability of settlements that, in many cases, is related to poorly building and the
612 absence of early warning systems.

613 The obtained results can be considered a relevant tool for future land planning in the town of Cuenca,
614 despite their resolution, limited to the municipalities area. The proposed method, using potentially
615 available data also for mid and low-income countries (i.e. landslide inventory and a proxy of
616 population distribution), has the potential to be applied in many contexts where a minimum dataset
617 is available or can be developed on the basis of either field or remote sensed data. The results of the
618 risk analysis are useful for ranking the municipalities in order of increasing risk and for supporting
619 decision-makers in prioritising funding for risk mitigation measures.

620

621 **Acknowledgements**

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