

24 **Abstract**

25 While radiative forcing and thus land surface temperatures have been shown to positively
26 correlate with fire severity, precipitation, and lightning strike frequency, the effects of human
27 activity on fire regimes remain difficult to disentangle from geophysical drivers given co-
28 variation between these factors. Here, I analyze fire regimes in the 1919-2012 period across
29 Canada and compare national trends to those of a latitudinal and elevational gradient in a region
30 experiencing exponentially increased anthropogenic activity in recent decades. Located along the
31 Canadian Rocky Mountains, the region is intended to serve as a proxy for future continental
32 conditions under current anthropogenic trajectories. Based on the findings, I argue that, for the
33 first time in millennia, fire regimes in the southern boreal zone have shifted on average from
34 large, lightning-caused fires to frequent, small, human-caused fires adjacent to human
35 transportation corridors. While warming is known to produce more severe fuel conditions,
36 human factors such as frequent fire ignitions, active fire suppression, industrial and recreational
37 activity, and forestry (i.e., stand aging) likely explain the reduction in mean fire size and annual
38 area burned. Here, I provide the first evidence of a southern boreal transition to Anthropocene
39 fire regimes without historical analogue, representing a dramatic departure from the conditions in
40 which these forests evolved. With ~28 Pg carbon stored in Canada's managed forests and
41 interspecific variation in albedo, these novel fire regimes carry direct implications for the Earth's
42 climate system.

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47 **Introduction**

48 The evolutionary history and paleorecord of North America’s boreal forests reflect
49 millennia of cold, dry, and fiery conditions (Gavin et al., 2007; He et al., 2012; Hu et al., 2006;
50 Kelly et al., 2013; Tinner et al., 2008). Global change in the Anthropocene (Crutzen and
51 Stoermer, 2000) has shifted each of these three conditions. Over the past half-century, boreal
52 forests warmed at twice the rate of the global mean (Intergovernmental Panel on Climate
53 Change, 2014). In southwestern Canada, recent climatic change produced warmer and wetter
54 conditions, significantly reduced snowfall, and related reductions in cryomass
55 (Intergovernmental Panel on Climate Change, 2014), or total mass of surface and ground water
56 in a frozen state. Warming is projected to accelerate in the near-term (Smith et al., 2015), with
57 the highest rates of warming expected to occur in mountainous regions (Miller, 2013) and higher
58 latitudes. The northernmost regions are experiencing the most severe temperature extremes of
59 the past 600 years through polar amplification (Miller, 2013; Tingley and Huybers, 2013).

60 The North American boreal is projected to migrate northward under warming, inducing a
61 net terrestrial loss of carbon storage (Koven, 2013; Scheffer et al., 2012). At lower elevations
62 and latitudes, extant tree species are expected to regenerate less frequently following disturbance
63 under warming, due to an increased frequency and magnitude of physiological drought
64 (Barichivich et al., 2014; Intergovernmental Panel on Climate Change, 2014; Nitschke et al.,
65 2010). Together, changes to regeneration and fire regimes may explain diminished recruitment
66 rates observed for Canada in recent years (Boisvert-Marsh et al., 2014; de Lafontaine and
67 Payette, 2011; Zhang et al., 2015). A reduction in area burned, without a compositional shift
68 toward deciduous trees, may further accelerate warming through reduced albedo (Amiro et al.,
69 2006).

70 Large stand-replacing fires have characterized circumpolar boreal forests for millennia,
71 reflected in the fire-resisting, -avoiding, and -embracing evolutionary strategies of the resident
72 tree species (Kelly et al., 2013; Rogers et al., 2015). Changes to fire regimes carry particular
73 importance in the North American boreal, where fire has been shown to regulate carbon flux
74 (Bond-Lamberty et al., 2007), energy partitioning (Amiro et al., 2006), compositional change,
75 and tree migration (de Lafontaine and Payette, 2011; Gavin et al., 2013). Warming has increased
76 the severity of fuel conditions in the boreal by increasing evaporative demand (Barichivich et al.,
77 2014; Intergovernmental Panel on Climate Change, 2014) and permafrost thaw (Baltzer et al.,
78 2014; Camill, 2005), accelerating carbon loss through an increased depth of ground-layer
79 burning, particularly for peatlands (Turetsky et al., 2011, 2015).

80 Recent burn rates for the North American boreal have been reported in excess of
81 Holocene (~11.7 kybp) fire regime limits (Kasischke and Turetsky, 2006; Kelly et al., 2013;
82 Marlon et al., 2013). The global area burned rapidly accelerated with the Industrial Revolution
83 before declining over the past century (Marlon et al., 2008, 2013). Unprecedented high burn rates
84 (short fire rotation periods) are evident for the Alaskan boreal in recent years (Kelly et al., 2013;
85 Turetsky et al., 2011). Yet, Alaska shows little agreement with other regions of the North
86 American boreal. The eastern Canadian boreal shows a fire frequency and biomass burning
87 maximum ~ 4.5 kybp and a steady decline thereafter, currently at a 7,000-year low, due to
88 decreased insolation, shorter fire seasons, and increased precipitation (Marlon et al., 2008, 2013).

89 More recently, early season warming has produced an increase in spring fire size, with
90 variation in fire patterns attributable to climate-related water table changes and post-glacial
91 topography (Ali et al., 2009, 2012). Regions of the western Canadian boreal similarly show
92 declines in area burned linked to increased precipitation over the past century (Meyn et al.,

93 2013). These studies indicate that co-varying patterns of solar radiation, temperature,
94 precipitation, physiological drought, and human activity explain global variability in the area
95 burned, with human activity playing an increasingly important role post-industrialization
96 (Marlon et al., 2008). The critical role of human activity is shown by a recent analysis of global
97 burned area (Andela et al., 2017). While short-term efficacy of fire suppression was shown for
98 Alberta (Cumming, 2005), long-term efficacy remains poorly understood.

99 In Scandinavia, boreal fire regimes shifted to their present state in the 17th century due to
100 increased human activity (Niklasson and Granström, 2000). In Niklasson & Granström (2002),
101 the fires-per-unit-area-time metric was used to indicate physical energetic constraints in the
102 configuration of fire regimes, based on fire frequency, size, and area burned per unit time,
103 following research on phase transitions in the classical Forest Fire Model (Drossel and Schwabl,
104 1992; Malamud et al., 1998). A recent analysis of global fire regimes supports the presence of
105 both physical energetic constraints and human-dominated fire regimes (Archibald et al., 2013).
106 Archibald *et al.* (2013) estimated energetic constraints from an expanded feature set that includes
107 fire frequency, size, intensity, season length, return interval, and area burned per unit time.

108 Similar to Niklasson & Granström (2000), Archibald *et al.* (2013) demonstrate that fire
109 frequency and size are inversely proportional for a given area burned per unit time. Fire
110 frequency strongly regulates fire intensity, while areas with shorter fire return intervals have
111 higher area burned per unit time. Longer fire seasons are related to higher human activity levels,
112 although difficult to uncouple from anthropogenic warming. Maximum fire size is characterized
113 by exponential decay and has a logarithmic relationship with area burned per unit time that
114 quickly approaches an asymptote (Archibald et al., 2013).

115 These findings reflect fundamental relationships between fire, climate, vegetation, and
116 human activity, supporting the theory of dual energetic controls (fuels and weather) on area
117 burned per unit time along productivity gradients (Archibald et al., 2013; Meyn et al., 2007,
118 2010). These studies also indicate that human activity poses a third fundamental energetic
119 constraint on fire regimes in the Anthropocene, alongside fuels and weather. Human activity may
120 explain recent changes to fire regimes in actively managed forests of southwestern Canada by
121 providing greater energetic inputs (ignitions), producing many small fires near human hotspots,
122 while reducing energy stores and spread potential (harvest, fuels management, and fire
123 suppression). These past-century changes to management are hypothesized to be evident in the
124 historical fire record.

125 Following pan-boreal (Bradshaw et al., 2009; Laurance et al., 2014) and regional trends
126 (Braid and Nielsen, 2015; Linke and McDermid, 2012), previous work has shown that increased
127 economic development in western Alberta expanded the road network into formerly remote
128 areas, facilitating increased access and use for economic and recreational purposes. Expanded
129 human activity is further evident in an increase in other linear features, such as oil and gas
130 pipelines, seismic lines, and power lines, as well as point features including one-hectare well-
131 sites (Linke and McDermid, 2012). While a number of studies have assessed disturbance patterns
132 here (Forest et al., 2008; Laberee et al., 2014; Nielsen et al., 2008), existing studies do not
133 explain the drivers of long-term disturbance variability critical to predicting future patterns in
134 simulation studies. Existing datasets may contain valuable information for discerning
135 relationships in space and time between human activity and fire, necessary for simulating
136 disturbance-related changes to understory solar irradiation. In the following sections, the effects
137 of past-century warming and increased human activity on fire regimes are assessed.

138 **Materials and methods**

139 Here, changes in the statistical patterns of historical wildfire data across western Alberta
140 and Canada are analyzed. The analysis focuses on climatic and anthropogenic changes to fire,
141 including variation in elevation, latitude, cause, size, frequency, and area burned along multiple
142 temporal resolutions, including annual, seasonal, monthly, and daily intervals. Fire seasons were
143 calculated as meteorological quarterly seasons. The analysis is structured to focus on proxies of
144 climatic change and human activity, based on known historical changes and the findings of
145 previous studies in the region. Although there exists significant variation in fire regimes across
146 Canada, national fire patterns provide a baseline for separating regional variation from overall
147 trends.

148 For the regional analysis, three data sources were used: the latest Canadian National Fire
149 Database (NFDB) fire perimeter data, NASA Shuttle RADAR Topography Mission (SRTM)
150 version 2 data processed using standard correction techniques (Reuter et al., 2007), and Natural
151 Regions and Subregions of Alberta for the biogeoclimatic zones (Natural Regions Committee,
152 2006). The data were subset to western Alberta and zonal statistics calculated for the minimum,
153 mean, and maximum elevation, as well as slope and aspect for each fire. The latitude and
154 longitude for each fire centroid was also calculated. The NFDB contains many relevant fire
155 attributes including the year, month, day, cause, source, and size. Using the year, month, and day
156 values, the ordinal date and season of fires were calculated. Using values for the elevation,
157 latitude, and ordinal date of each fire, foliar moisture content (FMC) was calculated for each fire.
158 To calculate FMC values, standard equations were applied from the Canadian Fire Behavior
159 Prediction (FBP) System (Hirsch, 1993).

160 Fire rotation period (FRP), or fire cycle, is a commonly applied metric to indicate the rate
161 of burning, with lower values indicating greater severity (Wagner, 1978). FRP is the average
162 time required for the sum of fire sizes within an area to equal the area in size, calculated over a
163 given time interval. FRP is often presented alongside the mean fire return interval (MFRI), the
164 average time interval between fires for a given area or site, as well as time-since-last-fire.

165

$$166 \quad FRP = \text{time interval} / (\text{sum of fire sizes burned in area} / \text{area size})$$

$$167 \quad MFRI = \text{time interval} / \text{number of fires in site or area}$$

168

169 Hence, FRP is the area-normalized MFRI. Applied to individual sites, FRP is equal to
170 MFRI. By normalizing for area, FRP provides more information about fire regimes at scales
171 greater than the individual site. MFRI values calculated for areas of different sizes are not
172 directly comparable, unless normalized for area, which yields FRP. FRP is applied in the
173 historical fire regime analysis. While other changes in the distribution of fires provide additional
174 information, FRP provides a single robust metric for fire regimes.

175 Fire size distributions were analyzed to detail variation in western Alberta and national
176 patterns, as well as changes to fire regimes between periods. This work follows a study on
177 lightning-caused fires in the boreal mixedwood region of Alberta, using the former LFDB
178 (Cumming, 2001) that showed that fire models should use a truncated exponential distribution to
179 prevent over-predicting large fires. Here, a Weibull distribution was fit to log-transformed fire
180 sizes. A right-tail Anderson-Darling maximum-goodness-of-fit estimation was used to adjust for
181 power-law behavior at the tail of the distribution. Hartigan's dip test was used to test for
182 bimodality. The expectation-maximization (EM) algorithm and Bayesian Monte Carlo Markov

183 Chain (MCMC) simulations were used to fit a mixed normal distribution. Changes to fire size
184 distributions were confirmed by Kullback-Leibler divergence and the Earth Mover's Distance
185 (EMD), or Wasserstein metric, commonly used for comparing empirical probability mass
186 functions (Gottschlich and Schuhmacher, 2014).

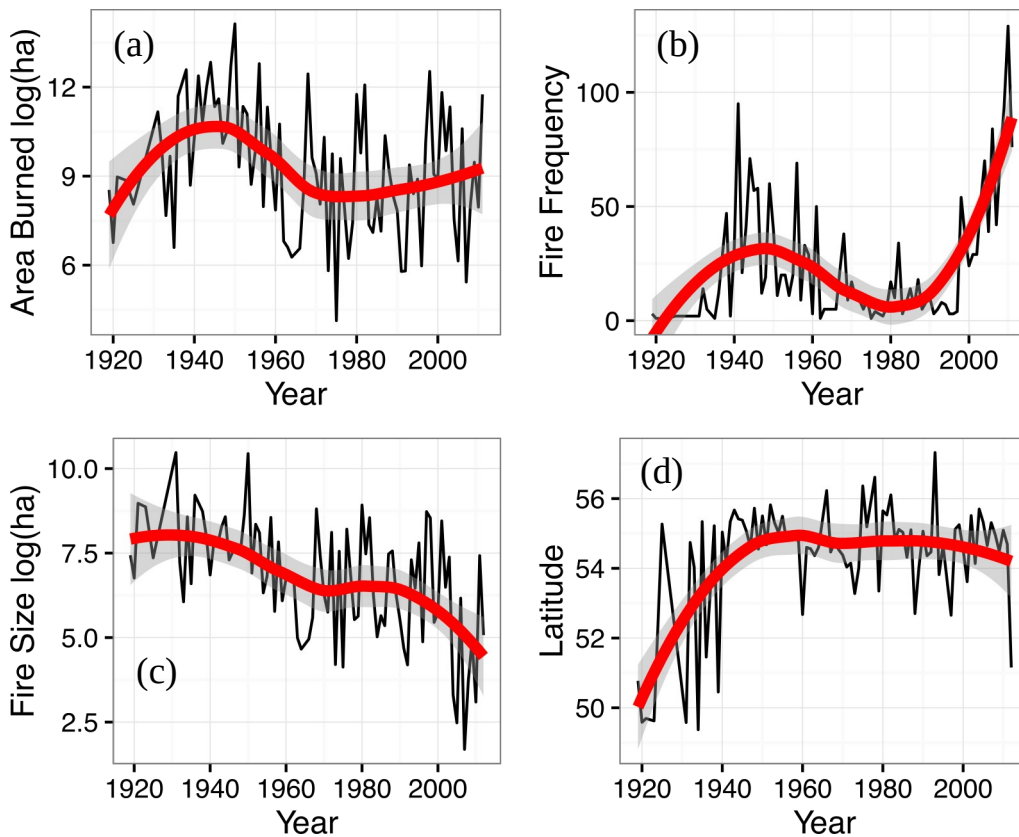
187 Fire regimes were temporally segmented using the binary segmentation algorithm (Scott
188 and Knott, 1974). While other change-point detection algorithms were tested, including pruned
189 exact linear time (Killick et al., 2011), e-divisive (Matteson and James, 2013), and e-divisive
190 with medians (James et al., 2014), binary segmentation showed optimal sensitivity to small
191 variations in the given task. Thus, binary segmentation was applied to classify fire regime
192 periods. First, fire regime periods were classified with *a priori* knowledge on changes to
193 management and climate. Periods of 30 years are used for compatibility with studies using 30-
194 year climate normal data. The four *a priori* fire regime periods are as follows: Pre-Suppression
195 (1923-1952); Early Suppression (1953-1982); Global Change (1983-2012); and, overlapping the
196 Global Change period, the Most Recent Decade (2003-2012). The Global Change period
197 corresponds to an acceleration of global change conditions (Steffen et al., 2007). The most recent
198 decade is included to represent recent trends independent of the three 30-year periods.

199 Software used to conduct this work includes ArcGIS 10.2 for spatial analysis, ENVI-IDL
200 5.2 for processing synthetic aperture RADAR data, R 3.1 for statistical analysis, and Python 2.7
201 for automation. The *seas* package for R was used for date-time conversion (Toews et al., 2007),
202 while the *fwi.fbp* package for R was used to calculate FMC values (Wang et al., 2014). The
203 *changepoint* (Killick and Eckley, 2014), *ecp* (James and Matteson, 2015), and
204 *BreakoutDetection* (James et al., 2014) packages for R were used to test change-point algorithms
205 for classifying fire regime periods.

206 **Results**

207 Across the full 90-year period in western Alberta, mean, maximum, and minimum fire
208 sizes declined. Fire frequency initially declined at an inflection point near 1950 before increasing
209 rapidly since approximately 1990. On average, over the 90-year period, fires declined in size by
210 142.6 ha per year, annual area burned declined by 3,450 ha per year, and fire frequency increased
211 by 5.44 fires per year (Figure 1).

212



213

214 **Figure 1. Mean annual trends for fires along western Alberta, 1919 to 2012:** (a) log of area burned; (b) fire
215 frequency; (c) log of fire size; (d) latitude in WGS84 (decimal degrees) coordinates; loess smoothing with 95%
216 confidence interval shown

217

218 In western Alberta, comprised predominantly of boreal forests, an inflection point in fire
219 regimes is apparent near 1970, with patterns in area burned, mean fire size, and mean fire
220 latitude changing thereafter; a rapid rise in fire frequency began a decade later (Figure 3.1). The
221 abruptness of the ~ 1970 and 1990 inflection points suggest a change in management, the former
222 potentially linked to an increase in oil and gas development in the boreal known to occur at the
223 time. An observed linear decrease in area burned and increase in fire frequency was independent
224 of elevation (high-elevation mean = -4607 ha/year, +3.3 fires/year; low-elevation mean = -29643
225 ha/year, +2.2 fires/year) and latitude (high-latitude mean = - 31211 ha/year, +3.0 fires/year; low-
226 latitude mean = - 3039 ha/year, +2.4 fires/year), based on median fire elevation and latitude. FRP
227 increased by 298% between the Pre-Suppression (1923-1952) and Global Change (1983-2012)
228 periods, indicating a three-fold reduction in fire regime severity during a period of warming
229 (Intergovernmental Panel on Climate Change, 2014).

230 FRP increased by 166% between the Pre-Suppression and Early Suppression (1953-1982)
231 periods, before increasing by another 50% between the Early Suppression and Global Change
232 periods. FRP in the Most Recent Decade (2003-2012) reflects patterns of the Global Change
233 period it overlaps, shorter by 0.1% at 923.9 years (Table 4.2). However, Most Recent Decade
234 fires were approximately twice as frequent and half the size of Global Change period fires
235 ($MFRI \Delta = -45\%$; annual *frequency* $\Delta = +82\%$; $MFS \Delta = -43.3\%$). The stability of FRP values
236 between the Global Change and Most Recent Decade periods is indicative of the temporal depth
237 of past-decade trends (Table 1).

238

239

240 **Table 1. Fire regime statistics by period for western Alberta;** mean fire return interval
 241 (MFRI) is shown in years for the full region rather than the mean site value, where
 242 $\text{Burned} = 1 / \text{MFRI} * \text{MFS} * \text{Years}$; $\text{FRP} = \text{Area} / \text{Burned} * \text{Years}$

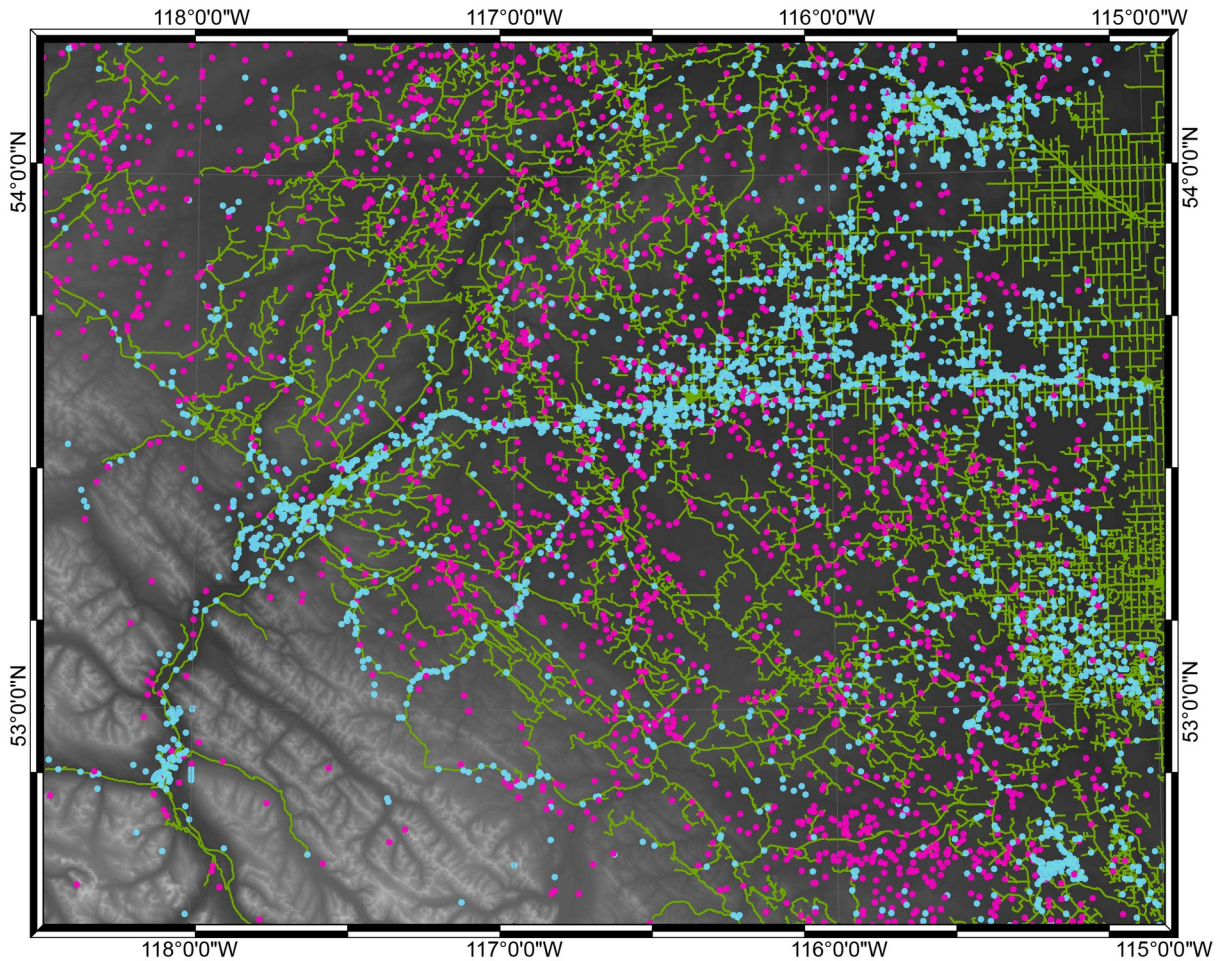
Period	Burned (ha)	Area (ha)	Fire Rotation Period (FRP, years)	Mean Fire Return Interval (MFRI, years)	Mean Fire Size (MFS, ha)
1923-1952	3,224,691	24,972,634	232.3	0.011	1,148.4
1953-1982	1,211,806	24,972,634	618.2	0.020	811.1
1983-2012	809,967	24,972,634	925.0	0.020	545.1
2003-2012	270,287	24,972,634	923.9	0.011	308.9

243
 244 Differences in the mean and variance of fire size between Early Suppression and Global
 245 Change periods were not statistically significant at a p -value threshold of 0.05 ($t = 1.69$, p -value
 246 $= 0.09$; $F = 1.19$, p -value $= 0.06$). Area burned declined substantially between these periods
 247 ($\text{Burned}_{ES} = 1,211,806$ ha, $\text{Burned}_{GC} = 809,967$ ha, $\Delta = -33.1\%$), even though remote monitoring
 248 improved. While MFRI remained stable across the Early Suppression and Global Change periods
 249 ($\text{MFRI}_{ES} = 0.0201$, $\text{MFRI}_{GC} = 0.0202$, $\Delta = +0.5\%$), mean fire size (MFS) declined at a rate
 250 equivalent to that of area burned ($\text{MFS}_{ES} = 811$ ha, $\text{MFS}_{GC} = 545$ ha, $\Delta = -33\%$). Thus, a decline
 251 in mean fire size best explains the reduction in area burned under warming in western Alberta.
 252 This is particularly evident in the Most Recent Decade, where FRP (area burned) was similar to
 253 the Global Change period it overlaps ($\Delta = -0.2\%$) as MFRI (fire frequency) increased by 81.8%.

254 In western Alberta, the ratio of human- to lightning-caused fires increased from 0.93:1 to
 255 1.39:1 (+33%) between the 1970s and 2000s. A spatial analysis of historical ignitions in western
 256 Alberta using NFDB point data demonstrates the proximity of small fires to areas of human
 257 activity, typically major roads and river valleys (fire distance from roads: mean = 2.2 km,

258 standard deviation = 4.8 km; fire distance from roads or surface water: mean = 297 m, standard
259 deviation = 363 m), supporting a human origin (Figure 2).

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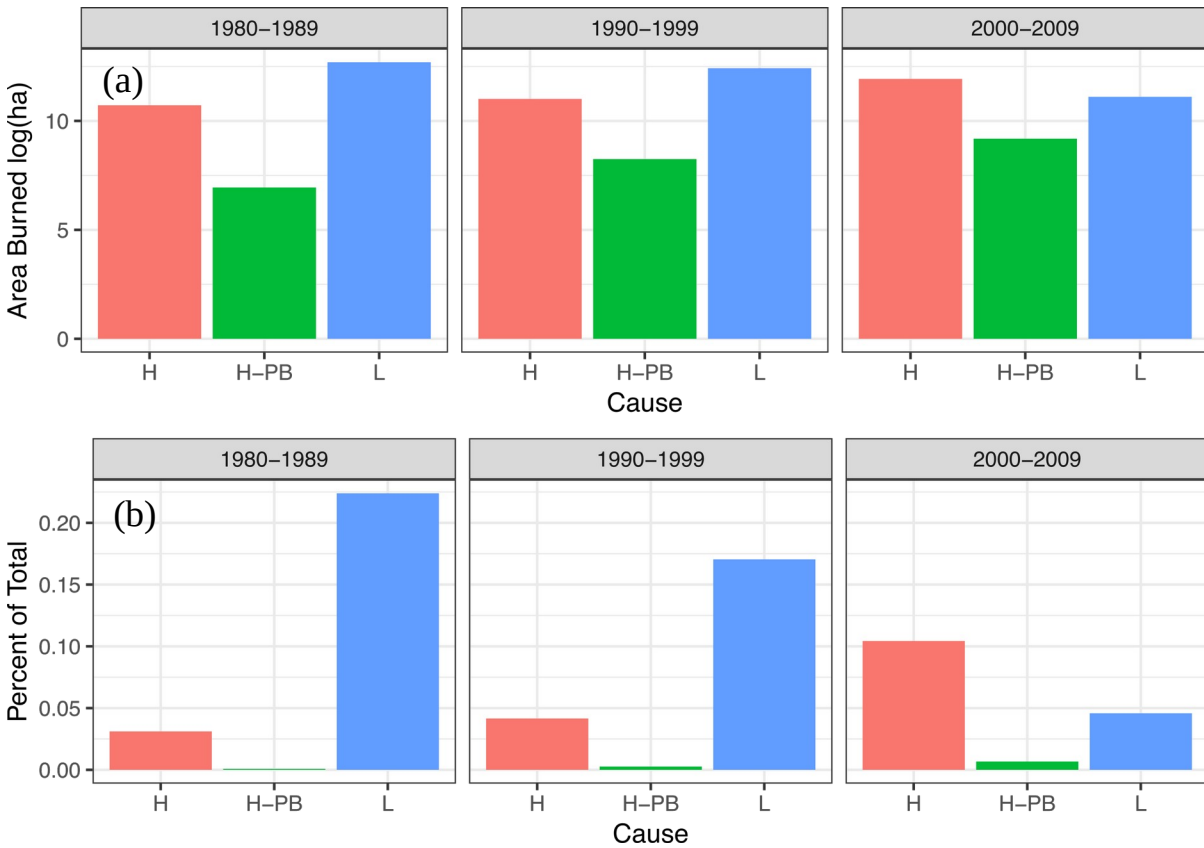
262 **Figure 2. Fire adjacency to roads by cause overlaid on SRTM 90 m elevation data in the vicinity of Hinton,**
263 **Alberta in NAD83 UTM 11N coordinates with WGS84 graticules: light blue = human-caused; magenta =**
264 **lightning-caused; green = roads; top = north**

265

266 Between the 1980s and 2000s, as Alberta's population doubled, the mean distance of
267 fires from roads declined by 40%, from 2.3 to 1.4 km. The mean distance of fires from roads or

268 surface water (rivers and lakes; proxies of human activity) declined by 32% across the same 30-
 269 year period, from 318 to 216 meters. Concurrently, annual fire frequency increased by 33%,
 270 from 6,035 to 9,054 fires, in the point data. The increasing influence of human activity in
 271 Alberta's fire regimes is apparent in the percentage of the total area burned attributable to
 272 sources over the past three decades (Figure 3).

273



274

275 **Figure 3. Decadal area burned by fire source for western Alberta;** (a) absolute area burned (ha); (b) percent of
 276 total area burned (ha); most fires were unknown in origin (not shown); H = human-caused; H-PB = prescribed burn;
 277 L = lightning

278

279 A decline in the relative influence of lightning on the total area burned in Alberta was
 280 offset by an increase in the percentage of area burned explained by human-caused fires. Between
 281 the 1970s and 2000s, the area burned increased by 34% in summer, fire frequency increased in
 282 spring and summer, and mean fire size increased by 83% in fall and 60% in spring for western
 283 Alberta (Tables 2a and 2b).
 284

Table 2a. Fire regime change by season

Season	Area	Number	Mean Size
Spring	-6.8%	+13.5%	-60%
Summer	+7.6%	+1%	+16.6%
Fall	-2.2%	-18.2%	+82.9%

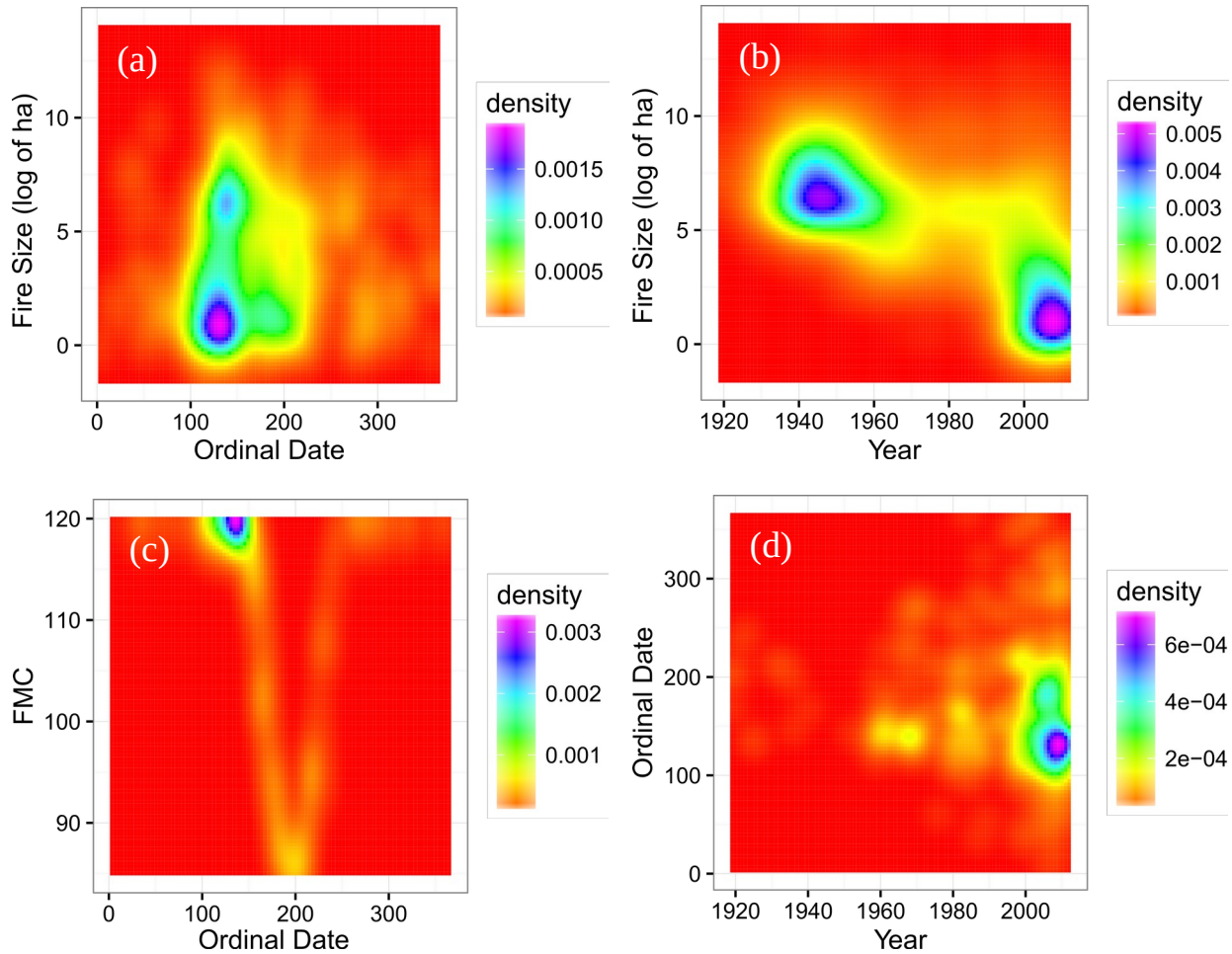
Table 2b. Current fire seasonality

Season	Area	Number	Mean Size
Spring	63.2%	51.2%	123%
Summer	33.8%	39.2%	86%
Fall	2.4%	6.8%	35%

285
 286 An analysis of fire seasonality related to the ‘spring dip’ in foliar moisture content (FMC)
 287 using standard formulations from the Canadian Fire Behavior Prediction System (Forestry
 288 Canada Fire Danger Group, 1992; Wotton et al., 2009), shows that the standard FMC equations
 289 are not suitable for western Alberta. Here, the modeled spring dip in FMC occurs approximately
 290 two months after the peak in fire frequency and size (Figures 4a and 4c) that likely corresponds
 291 to the true spring dip (Alexander and Cruz, 2013; Tymstra et al., 2007). The log of fire size
 292 shows the strongest density at 138 DOY (late April), followed by a second peak ~ 1 week later at
 293 a substantially larger fire size (Figure 4a). Meanwhile, modeled spring dip occurs at 200 DOY
 294 (Figure 4c). Across the full time period, fires declined in size following a structural change

295 around 1990 (Figure 4b). In recent years, fires were most frequent and concentrated earlier in the
296 season, with longer fire seasons (Figure 4d).

297



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299

300 **Figure 4. Two-dimensional kernel density estimation for fire frequency by ordinal date and year for western**
301 **Alberta, 1919 to 2012:** (a) log of mean fire size by ordinal date; (b); log of mean fire size by year (c) modeled foliar
302 moisture content (FMC) by ordinal date; (d) fire ordinal date by year

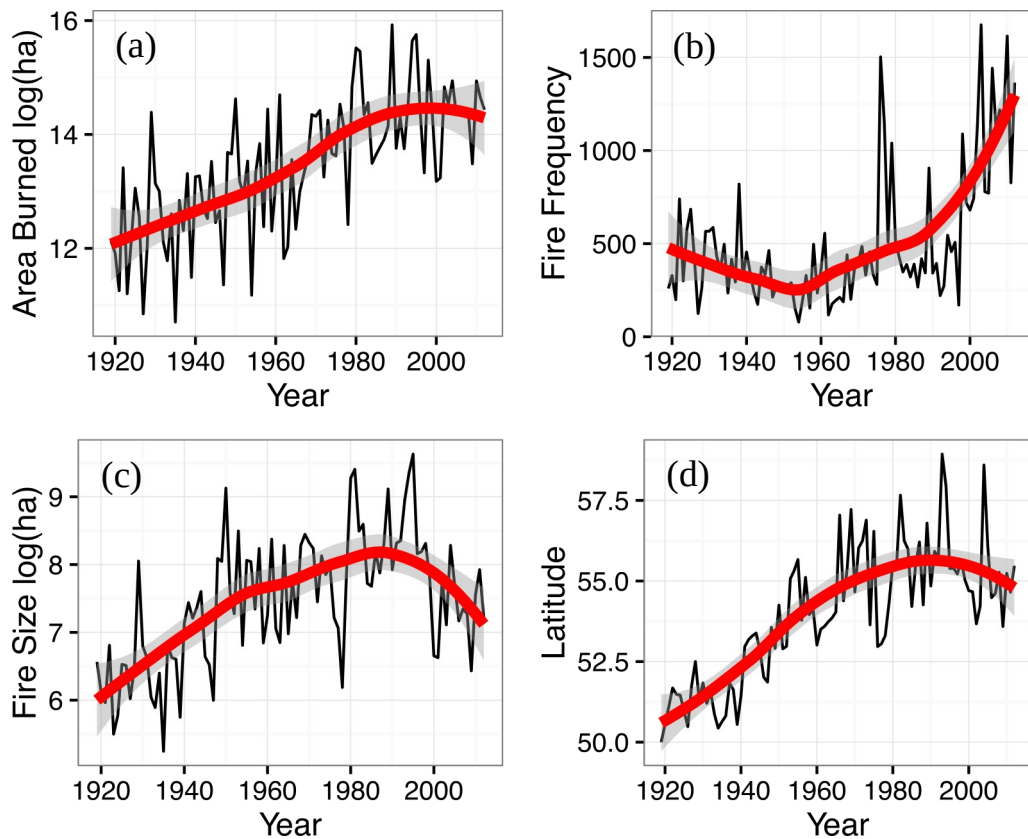
303

304 Across Canada, an increasing rate of area burned declined at a similar inflection point

305 around 1990, when mean fire size and latitude declined as fire frequency rapidly increased

306 (Figure 5). Between the 1990s and 2000s, fires nationwide declined in mean latitude at a rate of

307 14 km/year while fires in western Alberta declined at a rate of 24.4 km/year. Nationwide and in
 308 Alberta, lightning-caused fires decreased and human-caused fires increased in mean latitude
 309 during the period (nationwide = -6/+1.6 km/year; Alberta = -5.9/+1.5 km/year). Characteristic of
 310 the boreal, large lightning-caused fires ≥ 200 ha increased in mean latitude by 5.4 km/year
 311 nationwide and decreased by 10.1 km/year in Alberta. Since 1920, fires ≥ 200 ha shifted
 312 northward at a mean rate of 5.2 km/year ($R^2 = 0.13$; $p < 0.001$) nationally.
 313



314
 315 **Figure 5. Mean annual trends for fires Canada-wide, 1919 to 2012:** (a) log of area burned; (b) fire frequency; (c)
 316 log of fire size; (d) latitude in WGS84 (decimal degrees) coordinates; loess smoothing with 95% confidence interval
 317 shown; NFDB data prior to 1960 are known to be incomplete
 318

319 For the *a priori* classification in western Alberta, the Pre-suppression period (1923-1952)
320 is characterized by frequent fires and the largest annual area burned, while the Early Suppression
321 period (1953-1982) shows a sharp decrease in fire frequency and annual area burned, with the
322 lowest overall rates of each. The Global Change period (1983-2012) exhibits a rapid increase in
323 fire frequency but a relatively flat annual area burned. The Most Recent Decade (2003-2012),
324 shows the most rapid increase in fire frequency and the most rapid decline in mean fire size,
325 accompanied by a decline in area burned.

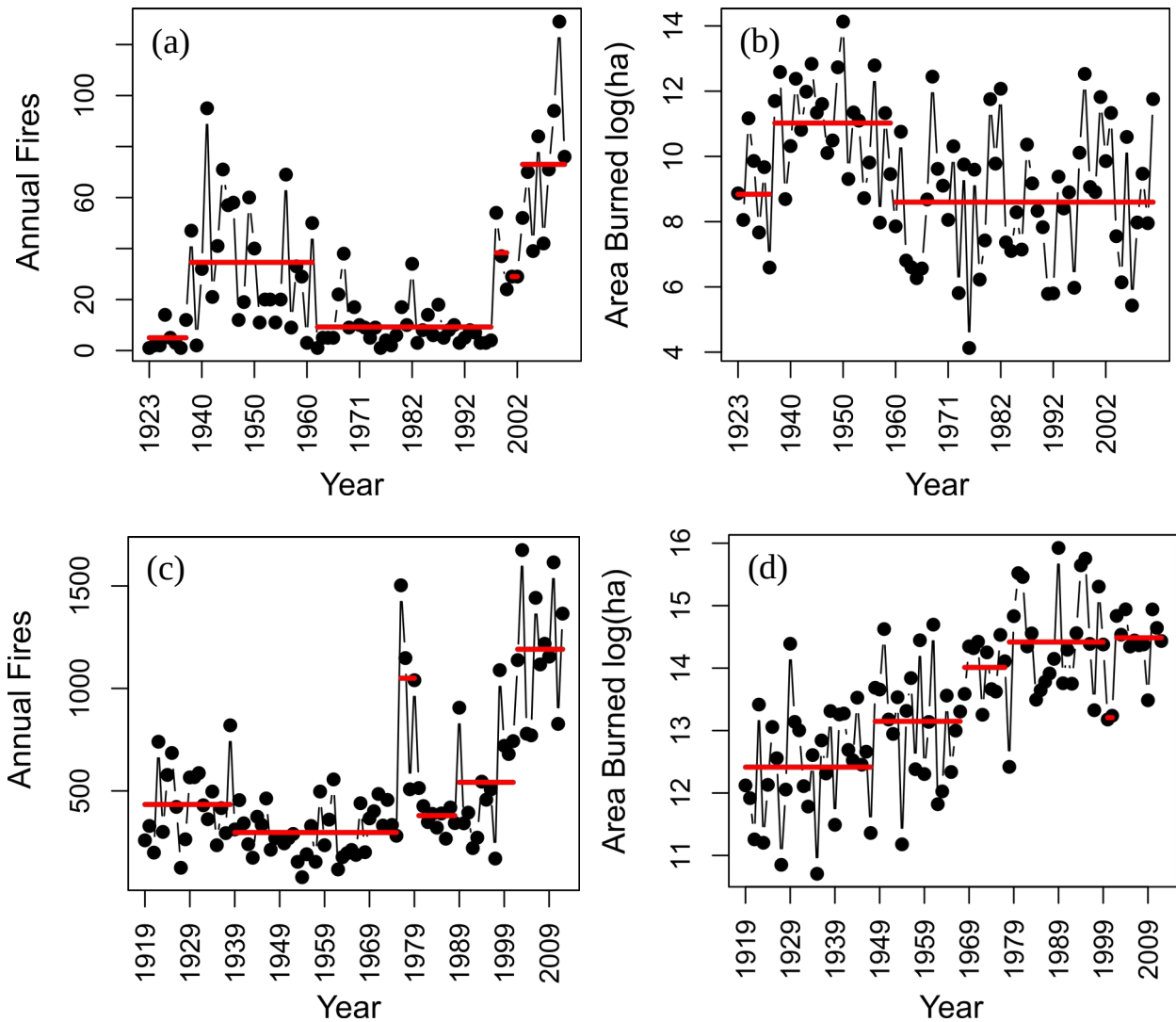
326 For western Alberta and Canada, fire regime period classification using the binary
327 segmentation change-point detection algorithm produced fire regime periods distinct from the *a*
328 *priori* classification. In Alberta, based on time-series of annual area burned and fire frequency
329 (Figures 6a and 6b), the algorithm shows an initial fire regime segmentation from the late 1930s
330 to the early 1960s, followed by another regime from the 1960s to the 1990s, and final regime
331 characterized by an increase in fire frequency from the 1990s to 2012. For Canada-wide fires, the
332 algorithm shows little consistency between fire regimes for the univariate annual area burned and
333 fire frequency time-series (Figures 6c and 6d). Nevertheless, the annual area burned time-series
334 shows approximate agreement with the *a priori* classification, with regime periods falling from
335 the early 1920s to ~ 1950, ~ 1950 to late 1960s, 1960s to late 1970s, and late 1970s to 2012.

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342 **Figure 6. Fire regime change-point segmentation using the binary segmentation algorithm:** (a) Alberta fire
 343 frequency by year; (b) Alberta total area burned by year; (c) Canada-wide fire frequency by year; (d) Canada-wide
 344 total area burned by year

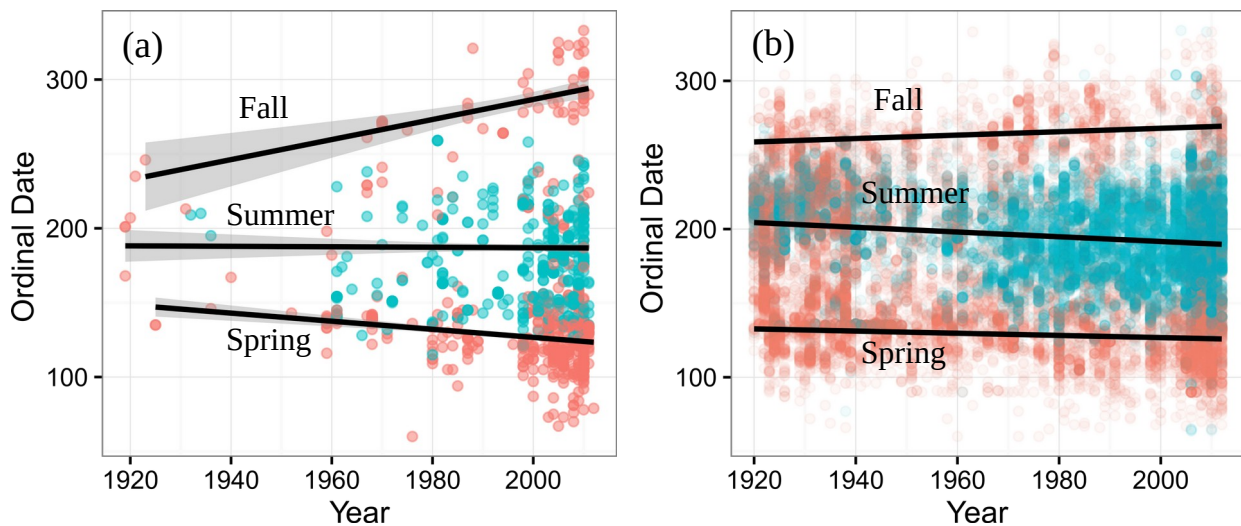
345

346 Fire regime periods differed for the two scales. Canada-wide, the 1940s through 1970s
 347 were characterized by infrequent fires and a steadily increasing area burned, while the inverse
 348 was true for Alberta. Nationwide, the first broad shift in fire regimes occurred during the 1970s
 349 with a spike in fire frequency and area burned. Yet, area burned was flat from ~ 1980. In western

350 Alberta, similar to nationwide patterns, fire frequency increased rapidly beginning ~ 1990. Yet,
351 Alberta showed little change in area burned from 1960 to 2012, despite strong variability within
352 the period (Figure 6).

353 Over the 90-year period, in western Alberta, fire seasons lengthened by ~ 60 days, or two
354 months (mean = +1.2 days/year), due to more frequent human-caused fires (mean = +9.2
355 fires/year) earlier and later in the season (Figure 7a). The fire season experienced a lower rate of
356 lengthening nationwide (Figure 7b). At both scales, lightning-caused fires were concentrated in
357 summer, while human-caused fires were concentrated in the spring and fall (Figure 7).

358



359

360 **Figure 7. Fire regime patterns nationwide and western Alberta changes in seasonality with linear models and**
361 **95% confidence intervals:** (a) western Alberta fire ordinal date by year and season; (b) Canada-wide ordinal date
362 by year, season, and cause; salmon = human; aqua = lightning; linear regression with 95% confidence interval
363 shown

364

365 Within western Alberta, the largest fire sizes and area burned occurred in the boreal,
366 followed by the foothills and Rocky Mountain regions. Within the boreal, the lowland

367 mixedwood subregions experienced a greater area burned than the highland subregions. Yet,
368 mean fire size and annual area burned declined in the boreal across the study period. Canada-
369 wide, the log-transformed fire size distribution for fires > 2 ha showed reasonable fit with a
370 Weibull distribution ($K-S = 0.02$; $\bar{\omega} = 1.60$; $A^2 = 27.34$; $AIC = 174679.6$; $BIC = 174696.8$), with
371 fit improving with the right-tail second-order Anderson-Darling (AD2R) statistic due to power-
372 law behavior at the tail. While the distribution of fire sizes nationwide showed unimodality per
373 Hartigan's dip test ($D = 0.003$, p -value = 0.11) despite visual evidence of bimodality, fire sizes in
374 western Alberta showed significant bimodality ($D = 0.02$, p -value = 0.002). Using a mixed
375 Gaussian model for western Alberta fire size, the two modes centered on μ of 1.2 and 6.2 log ha,
376 with Expectation-Maximization (EM) and Bayesian Markov Chain Monte Carlo (MCMC)
377 algorithms each converging to these values. This implies the presence of two dominant fire
378 regime phases in Alberta.

379 A further analysis reveals distinct changes in the fire size distribution over time. While
380 previous periods showed approximately Gaussian fire size distributions without skew, fires in the
381 Global Change period were strongly skewed toward smaller values. Bimodality of fire sizes for
382 all years in western Alberta is comprised of two distinct components: (1) frequent large fires in
383 1923-1952; (2) frequent small fires in 1983-2012. The Most Recent Decade showed the second
384 greatest K-L divergence ($D_{KL} = 0.84$, after the Pre-Suppression period ($D_{KL} = 1.05$), and greatest
385 distance from, the fire size distribution for all years, based on the Earth Mover's Distance (EMD)
386 or Wasserstein metric ($EMD = 5.15$). Fire regimes in Alberta thus reached a novel state in recent
387 years.

388 For Canada, monthly aggregations show mean fire size and total area burned were
389 typically largest in June, followed by July and May. Fire frequency peaked in July, followed by

390 June and May (Figures 8a – c). These findings are supported by daily resolution data. Given
391 increased temporal resolution, Gaussian and splines models indicate a typical fire frequency peak
392 between 184-185 DOY, mean fire size peak between 172-178 DOY, and area burned peak
393 between 171-178 DOY (Figures 8d – f). The splines models shows early season spikes in fire
394 frequency and areas burned corresponding with the ‘spring dip’ in foliar moisture content
395 indicated in Figure 3.4c – a sharp early season increase in the frequency and size of fires (Van
396 Wagner, 1967) – as well as a skewed fire frequency distribution. The log of mean daily fire size
397 centers at ~ 5.5 ha, while the log of mean daily area burned centers at ~ 6.5 ha. The log of daily
398 fire frequency shows a negative exponential distribution with a large λ value (Figures 8g – i).
399 This matches the typical model for the probability distribution of time-since-event for Poisson
400 processes, such as the probability of fire events, as in LANDIS-II (Sturtevant et al., 2009; Yang
401 et al., 2004).

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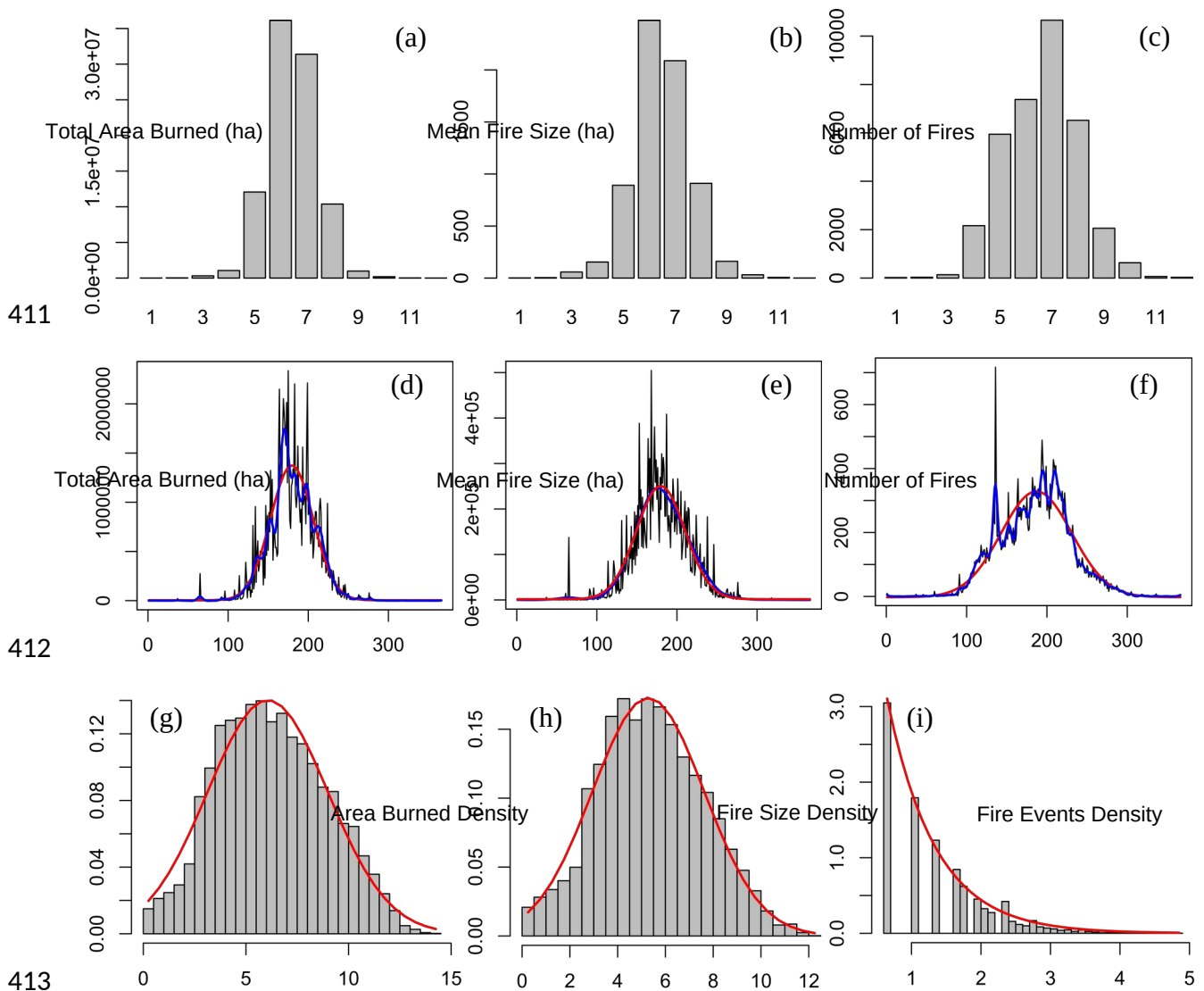
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414 **Figure 8. Monthly and daily patterns of fire frequency, mean fire size, and total area burned Canada-wide:**
 415 (a) total area burned by month; (b) mean fire size by month; (c) fire frequency by month; (d) total area burned by
 416 ordinal date; (e) mean fire size by ordinal date; (f) fire frequency by ordinal date; (g) log of daily area burned; (h)
 417 log of daily mean fire size; (i) log of daily fire frequency; recorded fire detection dates are used to calculate DOY
 418 values; blue and red lines in (d-f) are cubic splines and a Gaussian distribution fit, respectively, while the red line in
 419 (i) is an exponential distribution fit

420

421 **Discussion**

422 The distribution of fire sizes follows well-documented power-law behavior common to
423 self-organizing systems (Malamud et al., 1998; Reed and McKelvey, 2002), showing a heavy-
424 tailed distribution. Previous theoretical work suggested that fire size distributions should fit a
425 truncated Pareto distribution (Strauss et al., 1989). However, an empirical study of the boreal
426 mixedwood region of Alberta, using the former Large Fire Database for 1980-1998, showed
427 optimal model fit with a truncated exponential distribution (Cumming, 2001). The above results
428 suggest that the use of the AD2R goodness-of-fit statistic yields reasonable model fit with a
429 Weibull distribution for the logarithm of fire sizes.

430 The results illustrate that western Alberta experienced a sharp rise in human-caused fires
431 and area burned since 1990. This rise in human-caused fires likely combined with warming to
432 facilitate lengthening fire seasons in both early spring and late fall (spring mean = +0.26
433 days/year; fall mean = +0.67 days/year; $R^2 = 0.83$; $p < 0.001$), in agreement with previous
434 observations (Kasischke and Turetsky, 2006; Stocks et al., 2002). The combined lengthening of
435 fire seasons by 0.93 days/year is approximately five times faster than the Canada-wide average
436 of 0.2 days/year (spring mean = +0.07 days/year; fall mean = +0.13 days/year; $R^2 = 0.57$; $p <$
437 0.001). This difference in fire season lengthening rates is likely attributable to rapidly increasing
438 human activity in western Alberta, as well as data sparsity in early years. From the 1970s to the
439 2000s, human-caused fires accounted for a growing proportion of both annual fires (+9.8%) and
440 area burned (+38.9%) for fires of known cause.

441 While the majority of area burned continues to be produced by lightning-caused fires
442 nationally, this work observed a southern boreal shift to human-driven regimes characterized by
443 more frequent, smaller fires near human activity earlier and later in the year. Climatic warming

444 and a growing human presence are combining to create longer fire seasons, known to have
445 challenged managers in recent years (Tymstra et al., 2007). Fire frequency, area burned, and
446 mean fire size were greatest in spring for all regions, representing 51% of fires and 63% of area
447 burned, except the Rocky Mountain region, where fire frequency and size are greatest in summer
448 due to temperature constraints. The largest fires occurred in May, consistent with a ‘spring dip’
449 in foliar moisture content. Although this episodic decline in foliar moisture content remains
450 under investigation (Jolly et al., 2014), it is an important physiological phenomenon in these
451 forests (Alexander, 2010; Finney et al., 2013; Jolly et al., 2014; Little, 1970). Spring dip
452 typically corresponds to intense crown fire activity, producing the largest and most severe fires
453 of the fire season, which these data support.

454 Boreal fire regimes appear to be tracking a northward shift of boreal climatic conditions
455 (Koven, 2013), reducing the size and severity of fires in western Alberta, as southern boreal
456 ecosystems transition to Anthropocene fire regimes. Data from southeastern Canada indicate that
457 the in-migration of temperate species into the southeastern reaches of the American boreal is
458 already underway (Fisichelli et al., 2014). Fisichelli *et al.* (2014) proposes that the reduced size
459 of boreal fires, despite warming, is attributable to four key factors: (1) reduced surface fuel loads
460 from frequent small human-caused fires; (2) increased fire suppression; (3) reduced crown fuels
461 and/or forest fragmentation due to extractive industry activities; (4) a northward shift of boreal
462 climatic conditions, evidenced by changing wildfire patterns and climate-analogue vectors
463 (Koven, 2013).

464 A recent study shows demographic ageing for the region (Zhang et al., 2015), which may
465 further reduce surface fuels prior to gap formation and understory reinitiation. Previous studies
466 argue little effect of fire suppression on fire regimes in boreal and subalpine systems, as fuel

467 moisture shows greater importance than fuel load in models, while neither fire frequency nor
468 crown-fire potential were correlated with stand age (Johnson et al., 2001). Nevertheless, a shift
469 toward more frequent and smaller fires is evident for fire suppression regions (Kasischke and
470 Stocks, 2000). Subsequent analyses of Ontario and Alberta provide contrasting views on the
471 effectiveness of fire suppression in Canada (Bridge et al., 2005; Cumming, 2005).

472 The increasing extent and magnitude of industrial activity, recreational usage, and road
473 network expansion in formerly remote areas are combining with record temperature anomalies
474 (Kamae et al., 2014) to produce frequent ignitions and small fires around areas of human
475 activity. Harvest operations are widespread in these forests, reducing canopy fuels while
476 providing new ignition sources. A temporal lag of large fires following periodic pulses in pest
477 populations (Kurz et al., 2008) may amplify fuel conditions, fire regimes, and forest transition
478 rates. Increasingly warm and wet conditions may favor deciduous species in the southern boreal
479 (Terrier et al., 2012), producing a negative climatic feedback through increased summer albedo
480 (Amiro et al., 2006) while reducing the rate of fire spread (Dash et al., 2016).

481 An increase in fire suppression corresponds to an increasing human presence in
482 previously remote forested regions, related to a 619% increase in Alberta's population during the
483 1921-2011 period (Statistics Canada, 2011) and economic-related extractive industry activity
484 (Cross and Bowlby, 2006). The advent of fire suppression is indicated in the historical record by
485 reduced fire activity in the mid-20th century, following the 1950 Chinchaga wildfire in
486 northwestern BC and western Alberta, the largest recorded wildfire in North American history.
487 The disturbance legacy of this large fire is evident in the fire data, with few fires in its recovery
488 zone since, while surrounding boreal areas have burned frequently. This may partially explain
489 the observed decline in mean fire size, but it does not explain the accelerated decline in recent

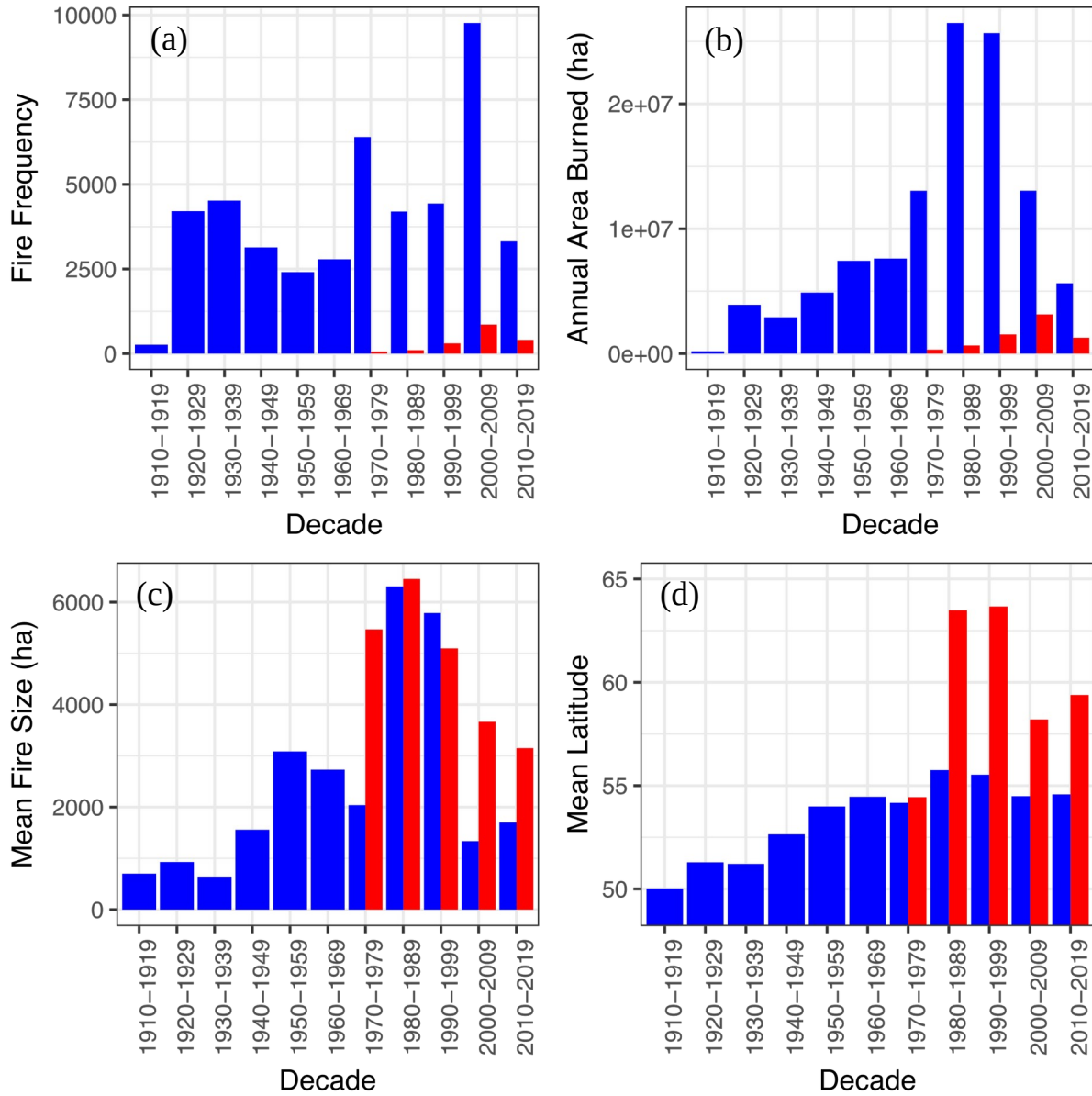
490 decades. The mean area burned by fires followed a similar trend, only rising in 1998 at the
491 beginning of an exponential-like increase in fire frequency, as described for other regions of the
492 boreal (Kasischke and Turetsky, 2006; Kelly et al., 2013). Research for Alberta, conducted
493 parallel to this work, selected a similar fire exclusion period start date of 1948, chosen for its
494 correspondence with the establishment of the Eastern Rockies Forest Conservation Board. This
495 work also shows a general lengthening of fire rotation periods compared to historical burn rates
496 (Rogean, 2016).

497 While one may infer that increased fire detection by satellites in recent decades (e.g.,
498 Landsat and MODIS) explains the observed rapid increase in fire frequency, decrease in fire size,
499 and increase in latitude of large lightning-caused fires during this period, an analysis of the
500 reported detection source rejects this hypothesis. Recent studies elsewhere in the boreal have
501 shown the effect of human activity on fire frequency (Gaglioti et al., 2016). While disturbance
502 detection source or instrument (spaceborne remote sensing versus traditional air and ground
503 methods) shows a statistically significant relationship for fire size ($p < 0.001$) and latitude ($p <$
504 0.001), it is not enough to explain recent fire regime changes.

505 Mean decadal fire frequency and area burned show little change due to inclusion of
506 spaceborne remote sensing over the past three decades (Figures 9a and 9b). Only mean fire
507 latitude and size were significantly impacted by detection source (Figures 9d and 9c), with the
508 effect greater for median values; an ANOVA indicates that latitude was more strongly affected
509 than fire size ($p = 7.39e-05$; $p < 2e-16$). Since the 1970s, spaceborne detection methods appear to
510 have substituted for traditional methods in northern regions. The mean decadal latitude of
511 lightning-caused fires > 200 ha peaked in the 1970s, prior to broad use of spaceborne

512 monitoring. Large lightning-caused fires were 2 degrees further north on average than fires
 513 200 ha, with a maximum of 5.3 degrees higher in the 1970s (WGS84 coordinates).

514



515

516 **Figure 9. Fire statistics by reported detection source Canada-wide:** (a) fire frequency by decade; (b) area burned
 517 by decade; (c) mean fire size by decade; (d) mean fire latitude by decade in WGS84 (decimal degrees) coordinates;
 518 blue = traditional detection source; red = modern remote sensing instruments

519 Thus, the contribution of spaceborne instruments to observed fire patterns remains small
520 relative to traditional methods. In the 2000s, spaceborne monitoring was used to detect less than
521 9% of recorded fires in Canada, despite reliable Landsat and MODIS coverage for the period
522 (Fensholt and Proud, 2012; Wulder et al., 2016). Even though spaceborne detection methods
523 produced a mean fire size twice that of traditional sources during the past decade, likely due to
524 the a combination of the coarse resolution of the MODIS hotspot product (Hantson et al., 2013)
525 and increased coverage in the north, the combined mean fire size sharply declined from 1990
526 onward. Furthermore, a rapid increase in the frequency of small human-caused fires in recent
527 decades may drive the mean fire latitude southward toward population centers. While this is
528 evident for fires of all sizes, large lightning-caused fires > 200 ha representative of classical
529 boreal fire regimes generally increased in latitude over the past 90 years, indicative of high-
530 latitude warming and an increased human presence in the north; disentangling these two factors,
531 as well as the inherent sampling bias of non-satellite detection methods, presents an opportunity
532 for future research. Nonetheless, a poleward shift of boreal fire regimes may correspond to a
533 northward migration of boreal forests under warming (D'Orangeville et al., 2016; Koven, 2013).
534 Further research leveraging the Landsat and/or AVHRR record is required to confirm this
535 dynamic.

536 While the inclusion of satellite disturbance detection data in recent decades should
537 increase the apparent area burned, the opposite is observed across regional and national scales.
538 At its peak prior to the current decade, in the 2000-2009 decade, spaceborne observations
539 represented 8% of fire observations and 19.4% of the total area burned; omitting these
540 observations leaves observed patterns generally intact. At its latitudinal peak in the 1990-1999

541 decade, spaceborne observations were 14% further northward on average compared to traditional
542 detection methods.

543 For western Alberta, where recent changes to fire regimes are greater than national
544 patterns, the detection source of fire observations shows no effect. According to the National Fire
545 Database, none of the fires in western Alberta were sourced from modern remote sensing
546 instruments. Hence, the analysis and related conclusions at the national and regional scales
547 remain valid. At the national scale, while the subtraction of remote sensing detected fires would
548 further increase FRP (reduce the rate of burning) during the Most Recent Decade, such large
549 fires were often drawn on a map by hand in previous decades. Absent additional information, I
550 estimate the historical fire size detection threshold at > 40 ha for western Alberta, based on
551 strong correspondence to a gamma fire size distribution at this threshold ($K-S = 0.028$ $\bar{\omega} = 0.231$;
552 $A^2 = 1.621$). Individual fire sizes between periods do not significantly differ in mean, but do
553 significantly differ in variance ($t = 0.750$, p -value = 0.454; $F = 301.180$, p -value $< 2.2e-16$).

554 Despite declines in fire size, latitude, and area burned for lightning-caused fires Canada-
555 wide between the 1990s and 2000s reported here (mean = -422 ha/year; mean = -15.6 km/year;
556 mean = -744,674 ha/year), a recent analysis of long-term warming suggests that these changes
557 are not climatic (Karl et al., 2015). Results for western Alberta show a regime shift toward
558 human-dominated fires in recent decades. By the 2000s, human-caused fires accounted for
559 58.1% of fires and 70.8% of area burned in western Alberta, surpassing the millennia-old
560 dominance of large lightning-caused fires. These findings contrast with Canada-wide changes
561 during the same period, where human-caused fires declined in contribution to the area burned
562 from 9% to 6%.

563 While human activity has long played a role in fire regimes in the boreal (Bowman et al.,
564 2011), Anthropocene conditions have recently combined to produce fire regimes without
565 historical analogue along the southern boreal. By analyzing fires > 200 ha before the 2000s, due
566 to limitations in the former national fire database, previous studies (Kasischke and Turetsky,
567 2006; Stocks et al., 2002) were unable to detect this regime shift. Fires < 200 ha in size represent
568 46.6% of fires in western Alberta (0.6% of area burned) and 59.3% of fires Canada-wide (0.9%
569 of area burned). Thus, while large fires continue to explain the area burned, they fail to explain
570 variation in fire frequency. As was shown, large recent changes to fire frequency are not
571 explained by the inclusion of spaceborne detection methods.

572 Our results for western Alberta contrast to previous studies suggesting that lightning
573 maintains a dominant role in annual area burned throughout the North American boreal
574 (Kasischke and Turetsky, 2006; Stocks et al., 2002). Here, more effective fire suppression
575 (Cumming, 2005) appears overwhelmed by a combination of warming and increased human
576 activity, beginning at an inflection point ~1970. At higher latitudes and elevation in Canada,
577 warming has been shown to increase biomass production (D'Orangeville et al., 2016; Hantson et
578 al., 2015), partially explaining an increased area burned here under the assumption of fuel
579 limitations.

580 An increased annual rate of fire frequency since 1980 corresponds with population
581 growth and increased economic activity in Alberta (Statistics Canada, 2011) combined with
582 rapid warming (Karl et al., 2015). Regional and national warming is evidenced by IPCC findings
583 (Intergovernmental Panel on Climate Change, 2014), previous fire regime analyses (Tymstra et
584 al., 2007; Wotton and Flannigan, 1993), and indirectly by aforementioned observed changes to
585 fire regimes Canada-wide. Human activity may explain most of the increase in the frequency of

586 small fires near roads and surface water, while warming also increases the frequency of lightning
587 strikes and severity of fire weather conditions (Krawchuk et al., 2009).

588 Although mean annual fire size and area burned declined in western Alberta over the past
589 decade, the effects of warming on burning appear to have been amplified, rather than attenuated,
590 by human activity. The data do not appear to support a previously reported non-linear U-shaped
591 relationship between human activity and the frequency of fire ignitions (Parisien et al., 2012;
592 Syphard et al., 2007). Due to the relative remoteness of Alberta's burnable land and small urban
593 areas (compared to populous regions, such as California), there appears to be an approximately
594 linear, rather than a U-shaped, distribution between fire frequency, area burned, and human
595 activity. Our results for western Alberta appear similar to findings for the Alaska boreal (Gaglioti
596 et al., 2016). Successful fire suppression efforts (Cumming, 2005) may partially account for the
597 decline in mean fire size nationally and in Alberta, as well as a declining national annual area
598 burned, despite warmer conditions with more frequent human-caused ignitions. High-frequency
599 small fires and extractive activities have likely also reduced forest fuels, which may together
600 explain an observed demographic shift in these forests (Zhang et al., 2015).

601 These patterns differ from other recent studies in the North American boreal including
602 Alaska (Kasischke and Turetsky, 2006; Stocks et al., 2002), which show a rapid rise in mean fire
603 size and annual area burned, based on analyses of previous historical fire database versions. The
604 results presented herein contradict both of these notions across regional and national scales,
605 showing greater agreement with paleoreconstructions from Alaska (Kelly et al., 2013), studies on
606 the relationship between human activity and fire frequency in the Alaskan boreal (Gaglioti et al.,
607 2016), and recent analyses indicating the presence of negative wildfire feedback mechanisms in
608 the North American boreal (Héon et al., 2014; Rogers et al., 2015).

609 Future studies should assess whether these trends are prominent across North America
610 and northern forests globally. A coupled climatic-human activity dynamic appears to explain the
611 observed changes in fire distribution. This is supported by a recent study showing a global
612 human-driven reduction in burned area (Andela et al., 2017). Studies should seek to better
613 delineate the causes of these patterns in terms of the precise roles of climatic, human, and forest
614 fuels mechanisms responsible. Of primary interest is the unexplained inflection point observed
615 around 1990, for both western Alberta and Canada, related to a rapid increase in fire frequency,
616 reduction in mean fire size, and reduction in area burned, despite warming. This poorly
617 understood inflection point appears to explain many observed dynamics. While historical
618 landcover and demographic change undoubtedly also play a critical role in explaining variations
619 in fire patterns, a dearth of detailed historical maps makes it difficult to assess, with remote
620 sensing records absent earlier than a few decades into the past. Future studies should investigate
621 the coupling of climatic change and human activity to better understand present and future
622 conditions, until more precise maps of landcover history are available.

623 Results indicate that the application of historical climate-fire correlations to general
624 circulation model projections, absent anthropogenic trajectories, carries diminished predictive
625 power in the Anthropocene. Short-term boreal ecological forecasts should include spatially
626 explicit dynamics of human-caused ignitions, fire suppression, and structural-demographic
627 changes to forest fuels related to increasing human activity. Long-term forecasts should further
628 include compositional change impacts on fuel conditions (Terrier et al., 2012), as well as coupled
629 climate feedbacks (Amiro et al., 2006).

630 These requirements may motivate the development of new terrestrial biosphere models
631 incorporating disturbance, succession, and energy partitioning processes, similar to recent hybrid

632 models (Bond-Lamberty et al., 2005; Scheller et al., 2007). The anthropogenically focused
633 Community Earth System Model (CESM1) revisions (Li et al., 2013) and Human-Earth System
634 Fire (HESFire) model (Le Page et al., 2015) represent such an approach, as do Eulerian grid-
635 based models such as WRF-Fire (Coen et al., 2012) and HIGRAD/FIRETEC (Colman and Linn,
636 2007).

637

638 ***Limitations***

639 This research relies on the best available fire history data for Canada (Burton et al., 2008;
640 Parisien et al., 2006; Stocks et al., 2002). Yet, the data contain known sampling biases toward
641 lower latitudes, larger fires of longer duration, and years subsequent to ~ 1960, particularly for
642 data on fire seasonality and cause. Thus, one would expect the data to show diminished mean fire
643 size and increased area burned, fire frequency, and mean fire latitude until the 1970s, when the
644 remote sensing record began with Landsat MSS. Yet, changes observed since the 1980s appear
645 robust to these sampling biases, given an increased satellite record. For improved estimates of
646 parameter uncertainty or model error, future studies may rely on hierarchical Bayesian modeling
647 with climate and anthropogenic data. While modern spaceborne imaging systems such as Planet
648 Doves (Hand, 2015) and deep learning techniques (LeCun et al., 2015) are poised to alleviate
649 sampling biases in historical fire maps over time by improving spatiotemporal resolution and
650 detection accuracy, the temporal depth of this remote sensing record remains limited, while the
651 substantial size of the data and neural networks remain cumbersome. Thus, early field
652 observations, airborne mapping, and paleorecords will remain indispensable for understanding
653 historical fire regimes.

654

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661 (<https://albertaparks.ca/albertaparksca/library/downloadable-data-sets/>), and NASA JPL
662 (<https://www2.jpl.nasa.gov/srtm/>).

663

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