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1 **The Influence of Grain Shape and Size on the Relationship** 2 **Between Porosity and Permeability in Sandstone**

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ABSTRACT

An accurate and reliable description of the relationship between porosity and permeability in geological materials is valuable in understanding subsurface fluid movement. This is of great importance for studies of reservoir characterisation, useful for energy exploitation, carbon capture, use and storage (CCUS) and groundwater contamination and remediation. Whilst the relationship between pore characteristics and porosity and permeability are well examined, there is scope for further investigation into the influence of grain characteristics on porosity and permeability due to the inherent relationship between grains and related pores. In this work we use digital image analysis (DIA) of reconstructed 3D X-ray micro computed tomographic (μ CT) images to measure porosity, permeability and segment individual grains enabling the measurement of grain shape (sphericity) and size (Ferret diameter). We compare two marker-based watershed workflows to grain boundary segmentation before applying the most reliable one to our images. We found there to be a positive relationship between grain sphericity and porosity according to $\phi = 1.22\phi_s - 0.42$ whereas no such relationship exists with grain size. We applied our grain shape and size measurements to calculate a Kozeny-Carman (K-C) porosity-permeability fit which was found to be unsatisfactory, possibly due to significant deviation from the K-C assumption that grains are spherical. Therefore, we show that a simpler fit of the form $K = 10^{5.54} \phi^{3.7}$, excluding any influence of grain characteristics, is most

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43 suitable for the studied materials and that grain shape and size is not influential
44 on the porosity-permeability relationship in a K-C paradigm.

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INTRODUCTION

47 The relationship between porosity and permeability is very significant for
48 reservoir characterisation studies applied to energy exploitation, carbon storage
49 and aquifer contamination and remediation. Constraining the relationship
50 between these two important reservoir parameters is beneficial as measurement
51 of porosity alone can then be used to predict permeability, which is typically
52 expensive and time consuming to measure both physically in a lab and
53 computationally using digital image analysis (DIA). Furthermore, permeability can
54 only be measured directly in the lab on small scale samples or in the field at the
55 macro scale using pump tests, producing two results which often do not closely
56 agree. Therefore, identification of a reliable and accurate relationship between
57 porosity and permeability using computed tomography (CT) imaging could have
58 far-reaching implications for reconciling this issue.

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Modelling a Porosity-Permeability Relationship

61 The Kozeny-Carman (K-C) relationship, proposed by Kozeny (1927) and later
62 modified by Carman (1937), is a simple yet broadly effective and widely used
63 (Mavko and Nur 1997; de Lima and Sri Niwas 2000; Urumovic and Urumovic Sr.

64 2014; Berg 2014; Hommel et al. 2018) technique of relating porosity to
65 permeability. Bear (1972) suggested a modification to the K-C equation which
66 allows grain diameter to be employed as a component which influences the
67 permeability. Additionally, Hommel et al. (2018) show that an additional grain
68 sphericity term may also be used. Whilst a K-C-based approach is successful in
69 many instances, its accuracy may be questioned when applied to materials which
70 possess a significant proportion of grains which deviate substantially from being
71 spherical. The limitation of a K-C approach is that grains are considered spherical
72 and packed in a regular arrangement; allowing pores to be considered as capillary
73 bundles. The inherent relationship between the pore structure and the grains
74 which create the pore space indicates that a detailed investigation of grain
75 characteristics is of utmost importance in understanding the porosity-
76 permeability relationship.

77 In this work we aim to investigate whether the inclusion of grain sphericity
78 and 3D Feret diameter (referred to herein as grain size) in a K-C paradigm
79 facilitates a better quality fit to the relationship between porosity and
80 permeability. We compare our modified K-C approach to a simpler fit using
81 porosity and permeability measurements alone, excluding any influence of grain
82 shape or size. To do so, the individual relationships between porosity and
83 permeability and grain sphericity and size are investigated and considered in light
84 of the concept of grain anisotropy, as introduced by Nabawy (2014).

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A Methodology for Making Digital 3D Grain Measurements

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Whilst grain size and shape measurement has traditionally been done manually using callipers and sieve analysis (W. D. Keller 1945; Schäfer and Teyssen 1987; Wang et al. 2013; Suhr et al. 2018) we have used digital image analysis (DIA) to segment individual grains in 3D using reconstructed X-ray micro computed tomographic (μ CT) image stacks of each sample. μ CT imaging has been used in a wide variety of fields related to geosciences since its rise in popularity as a non-destructive and high resolution image acquisition technique (Blunt et al. 2013; Bultreys et al. 2015; Thomson et al. 2018, 2020b; Payton et al. 2021). When paired with DIA, large amounts of quantitative and visually useful data may be obtained. Unlike when using optical imaging, X-ray imaging is dependent primarily on phase density therefore, grain boundaries are difficult to identify, particularly in a tightly packed sandstone.

In this work we discuss and investigate grain segmentation using two relatively simple marker-based watershed workflows. Watershed algorithms, established by Beucher & Meyer (2018), split a phase up into individual components by treating the image as a topographic surface, identifying topographic lows and assigning a seed point to each. Flooding from each seed point allows digital watersheds to be identified and are used to define the boundaries between individual features (Sun et al. 2019). The challenge arises

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106 from making correct identification of marker points so as not to have multiple
107 grains sharing one marker (undersegmentation) or the opposite where multiple
108 markers are assigned to a single grain (oversegmentation). Techniques such as
109 the bring up (Kong and Fonseca 2018; Leonti et al. 2020) and bring down (Shi and
110 Yan 2015; Sun et al. 2019) methods have been developed to try and tackle this
111 issue but can often be computationally demanding and may still produce
112 inaccuracies.

113 Segmentation of the solid phase alone allows identification of individual
114 grains which can then be measured digitally in 3D. Segmentation is arguably the
115 most important and usually most difficult process in DIA (Campbell et al. 2018)
116 given that poor segmentation will directly result in poor and likely misleading
117 results. It is notoriously difficult to segment features within a given phase which
118 are touching, consequently many techniques have been developed to tackle this
119 challenge, often providing unique solutions to a given sample set or type of
120 sample (shelly, angular, rounded, etc...) (Campbell et al. 2018; Kong and Fonseca
121 2018; Furat et al. 2019; Leonti et al. 2020) as there is not a one size fits all solution
122 (Campbell et al. 2018).

123 We assess two segmentation workflows and use the most effective to
124 analyse a collection of 22 sandstone samples from three different geological
125 formations (i.e., Wilmslow Sandstone Formation, Sellafeld, UK; Brae Formation
126 Sandstone, Miller Field, North Sea, UK; Minard Formation Sandstone, Porcupine

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127 Basin, North Atlantic Ocean). Finally, we use the grain measurements alongside
128 digital measurements of porosity and permeability to investigate the quality of a
129 K-C-based fit to the porosity-permeability relationship using grain shape and size
130 inputs.

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METHODS

133 A variety of sandstone samples have been selected from several different
134 reservoir units which host significant levels of porosity. Samples from the
135 Wilmslow Sandstone Formation (Sellafield, UK; Payton et al. 2021), Brae Formation
136 Sandstone (North Sea, UK; Thomson et al. 2020b) and the Porcupine Basin (North
137 Atlantic Ocean) were acquired and imaged at the London Natural History Museum
138 Imaging and Analysis Centre. Table 1 summarises the materials used in this work
139 and specifies the associated literature detailing initial sample imaging where
140 relevant. We chose to exclude samples which exhibited no connected porosity
141 and therefore no permeability for the purpose of this study.

142 The material pertaining to the Porcupine Basin was collected and prepared
143 using the same technique outlined by Thomson et al. (2020b) and Payton et al.
144 (2021). From each sample a mini plug measuring 5 mm in diameter and 10 mm in
145 length was cut and imaged using X-ray micro computed tomography (μ CT),
146 detailed by Payton et al. (2021). For further information about the voxel size and

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147 subsampled volume of each sample we refer the reader to the Supplementary
148 Information.

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150 *Image Processing*

151 The acquired μ CT image stacks of each sample underwent pre-processing
152 using the commercial software package PerGeos (v1.7.0). From each image stack
153 a sub-volume was extracted to remove external voxels and any image slices which
154 contained significant beam hardening artefacts. In order to aid the segmentation
155 process we employed a non-local means filter which enhances the contrast
156 between greyscale phases and removes speckled noise throughout the images
157 (Buades et al. 2008, 2010).

158

159 *Porosity and Permeability*

160 We followed the method detailed by Payton et al. (2021) to measure
161 porosity and permeability - a brief outline is described here. We made use of the
162 well-known automatic binary segmentation algorithm designed by Otsu (1979) to
163 separate and label the solid grain phase and pore space. In some cases, it was
164 necessary to constrain the greyscale range over which the algorithm was allowed
165 to operate on where exceedingly bright phases were present which meant darker
166 grains and darker pore space were not automatically separated.

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186 sticks to represent throats. We created PNM of the connected porosity following
187 the methodology detailed in Payton et al. (2021) and references therein. Each
188 PNM may be interrogated to provide information about each pore including
189 radius and coordination number, and each throat including radius and length.

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Grain Segmentation

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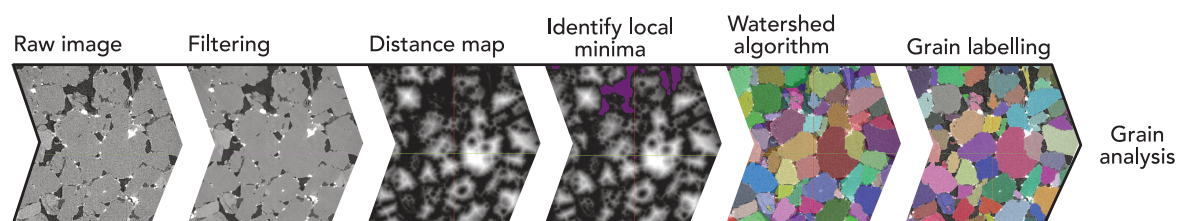
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Segmentation of individual features in μ CT images has traditionally been performed using the marker-based watershed approach detailed by Beucher & Meyer (2018). This technique has been widely used in a variety of fields (Barraud 2006; Cristoforetti et al. 2008; Veta et al. 2011; Huang et al. 2018; Xue et al. 2021) to identify and split individual features in digital images. The general steps in using a watershed algorithm are shown in Figure 1 (for a more detailed description of how a watershed algorithm operates we refer the reader to Kong & Fonseca (2018) and Sun et al. (2019)). We chose to follow the workflow of watershed segmentation of grains described by Fei & Narsilio (2020) which is shown to be

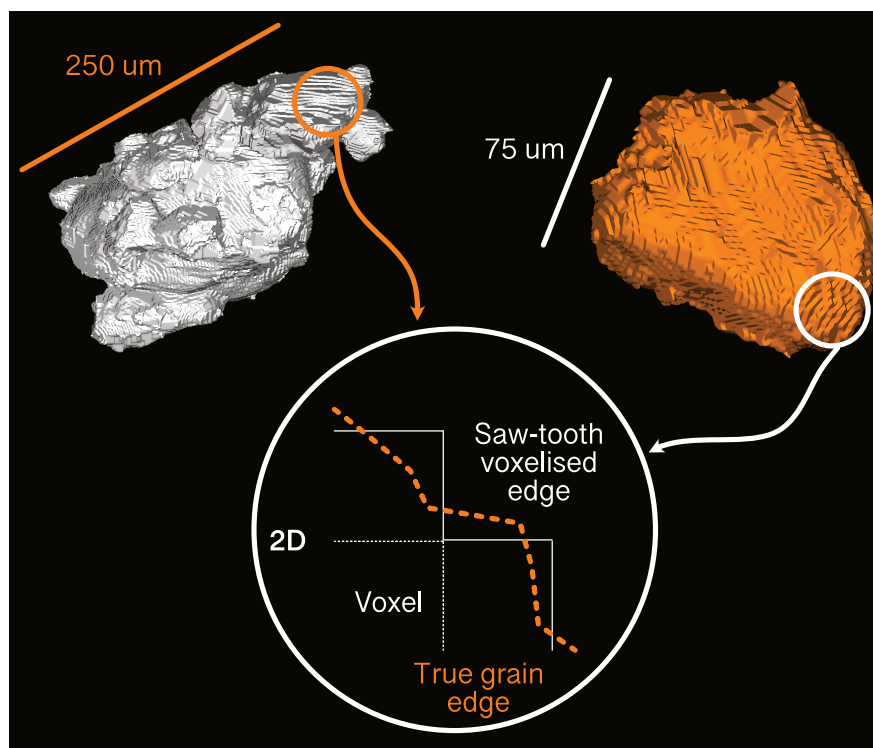


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successful in separating grains in a variety of different sand samples which bare
202 some resemblance to the materials investigated here.

203

Figure 1



221 **Figure 2**

222 Whilst smoothing algorithms can be applied to reduce this effect,
223 determining appropriate parameters for such algorithms becomes heavily
224 subjective and can cause undesirable deformation of the individual grains such as
225 volume loss. Moreover, using the same degree of smoothing on a very small and
226 a very large grain will have different impacts on the resulting shape. Consequently,
227 we chose to omit the use of any smoothing tools prior to our measurements being
228 made.

229 The automated nature of the MorphoLibJ and 3D Suite plug-ins enables this
230 analysis to be carried out simply as well as rapidly with low computational cost.
231 Sphericity is measured between 0 and 1 where 1 represents a perfect sphere. We
232 used Feret diameter as the representative grain size for all statistical analyses in

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233 this work. Some of the grain size analyses performed use phi (ϕ) units, calculated
234 from grain size values in millimetres according to:

$$\phi = -\log_2 D \quad (3)$$

235 where D is the grain diameter. We calculated the graphic mean grain size (M_z)
236 after Folk (1980) according to the following formula:

$$M_z = \frac{(\phi_{16} + \phi_{50} + \phi_{84})}{3}, \quad (4)$$

237 where ϕ_{84} represents the ϕ value at the 84th percentile. We calculated the
238 'inclusive graphic standard deviation' introduced by Folk (1980) to determine the
239 sorting (ϕ_1) of our samples using the following formula:

$$\phi_1 = \frac{\phi_{84} - \phi_{16}}{4} + \frac{\phi_{95} - \phi_5}{6.6}. \quad (5)$$

240 We then classified the sorting of our samples following the accompanying scheme
241 defined by Folk (1980) where a smaller ϕ_1 value is representative of better sorting.

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RESULTS

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Application of the Proposed Methodology

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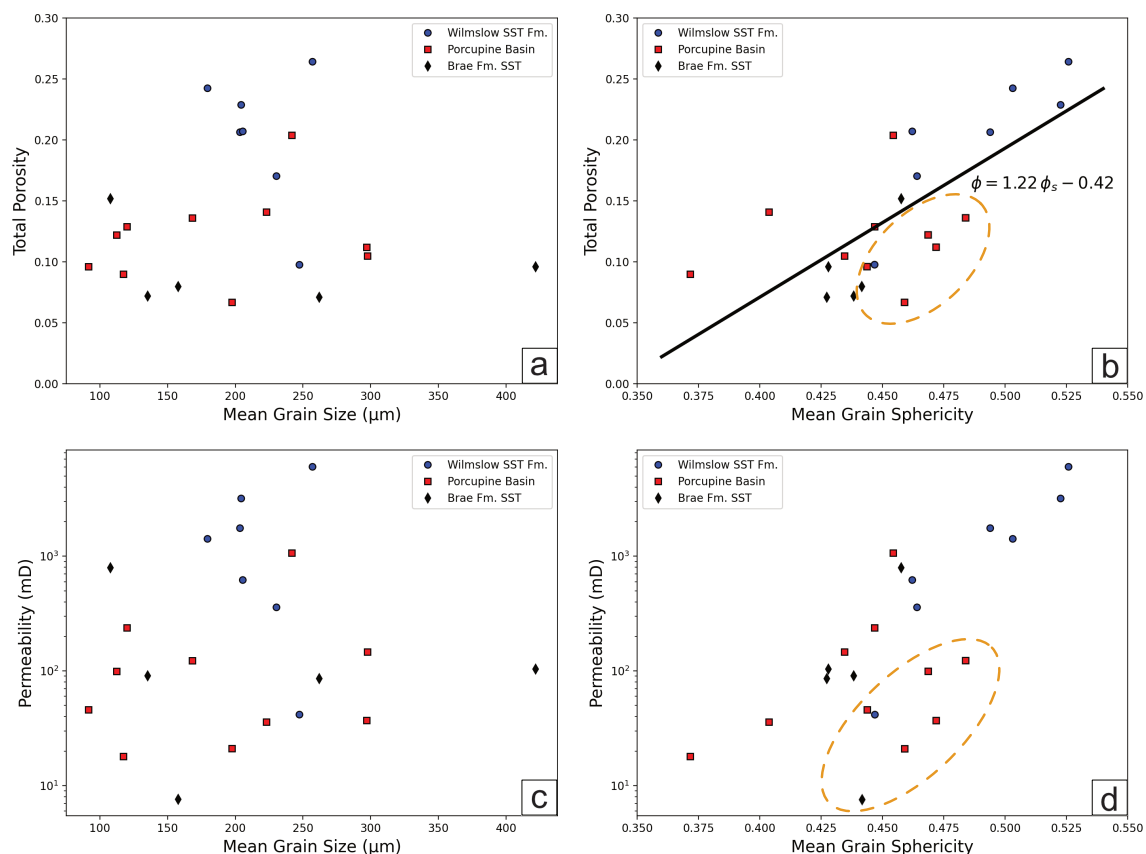
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Each study sample was analysed in terms of grain characteristics and the results are reported in Table 2. The accompanying porosity and permeability results are reported in Table 3, measured in this article and by Thomson et al. (2020b) and Payton et al. (2021). Figure 3 shows the relationships among mean grain size, mean grain sphericity, porosity, and permeability. No clear relationship between grain size and sample porosity or permeability is observed (Figs. 3a and

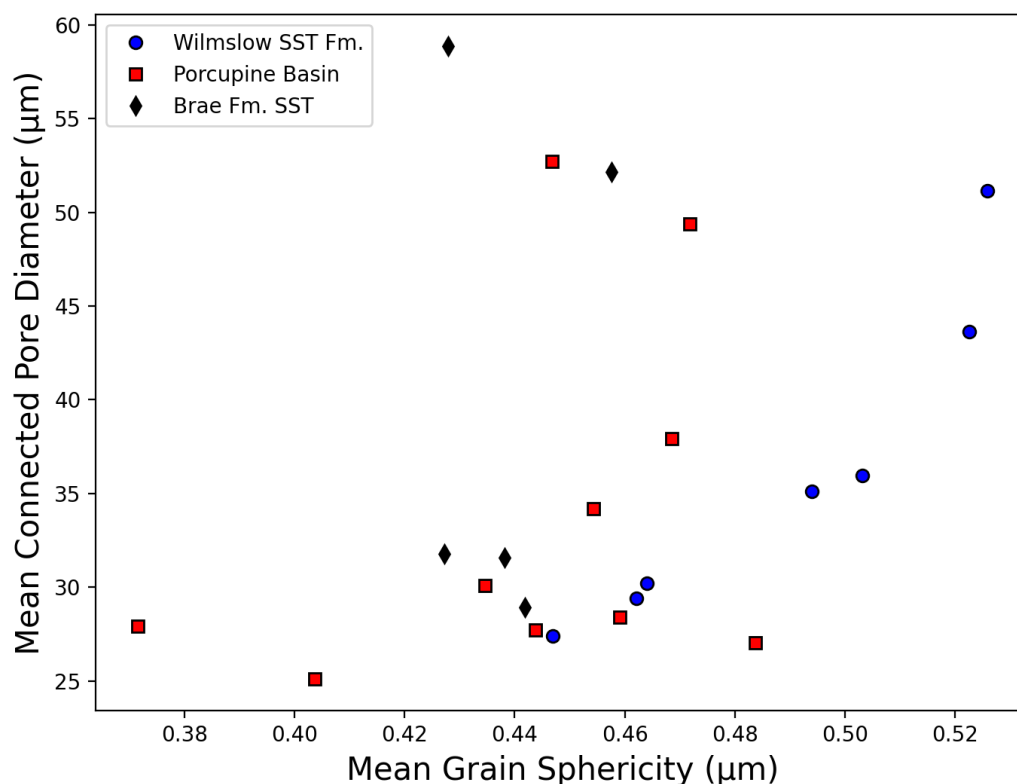
251 3c). Despite this, we see a much clearer positive correlation between the grain
 252 sphericity and porosity and permeability (Figs. 3b and 3d). This suggests that the
 253 shape or anisotropy of the grains has a direct influence on the pore structure
 254 whereas the size of the grains does not. Figure 3d highlights a collection of seven



255 outliers showing the same relationship but offset from the dominant trend
 256 between mean grain sphericity and permeability. The same collection of data
 257 points is highlighted in Figure 3b, plotting mean grain sphericity against total
 258 porosity, where they are not obviously misaligned with the rest of the data points.
 259 This indicates that these apparent outliers, in the case of permeability, result from
 260 a characteristic of the sample which is independent of porosity but not
 261 permeability.

262 **Figure 3**

263 As the intergranular porosity is fundamentally governed by the grains
264 themselves, we investigated the relationship between the pore structure and
265 grain sphericity. Figure 4 shows a generally positive relationship between grain
266 sphericity and the connected pore diameter except for four apparent outliers
267 across all three sample suites. Of these four outliers, two belong to the group of
268 seven identified in Figure 3d and two do not. However, the cause for the
269 occurrence of these four outliers is unclear and it seems that there is no
270 correlation between these four outliers and other measured factors such as
271 sorting and grain size.



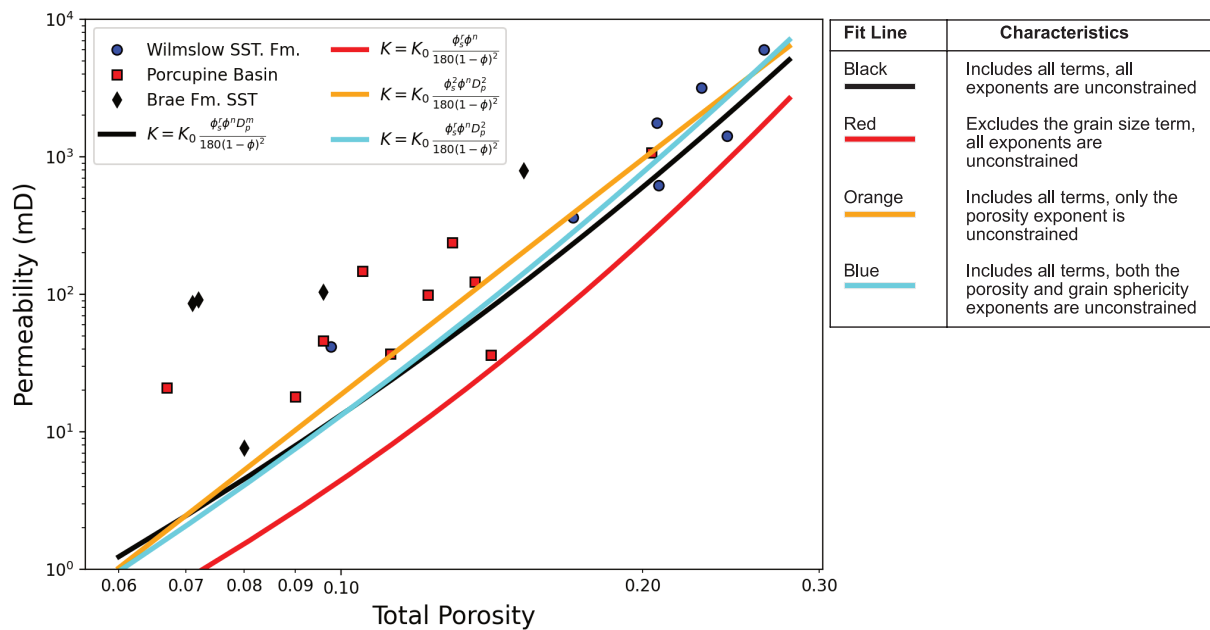
272 **Figure 4**

273 *Impact of Grain Shape on the Porosity-Permeability Relationship*

274 Our results show that the shape of the grains in a sample has an impact on
 275 the porosity and permeability. Therefore, it is reasonable to assume that the
 276 porosity-permeability relationship could be better constrained through
 277 incorporating the grain shape into the fit equation. Accordingly, we employed a
 278 modified Kozeny-Carman equation discussed by Hommel et al. (2018),

$$K = K_0 \frac{\phi_s^r \phi^n D_p^m}{180 (1 - \phi)^2}, \quad (6)$$

279 which incorporates the grain sphericity, ϕ_s and size, D_p alongside porosity, ϕ and
 280 a permeability constant, K_0 to calculate a porosity-permeability fit. We imposed a
 281 variety of constraints on the fit with regards to the three constant exponents: n ,
 282 m and r , applicable to porosity, grain size and grain sphericity respectively (Fig. 5),
 283 to determine the best fit with the lowest root mean square error (RMSE).



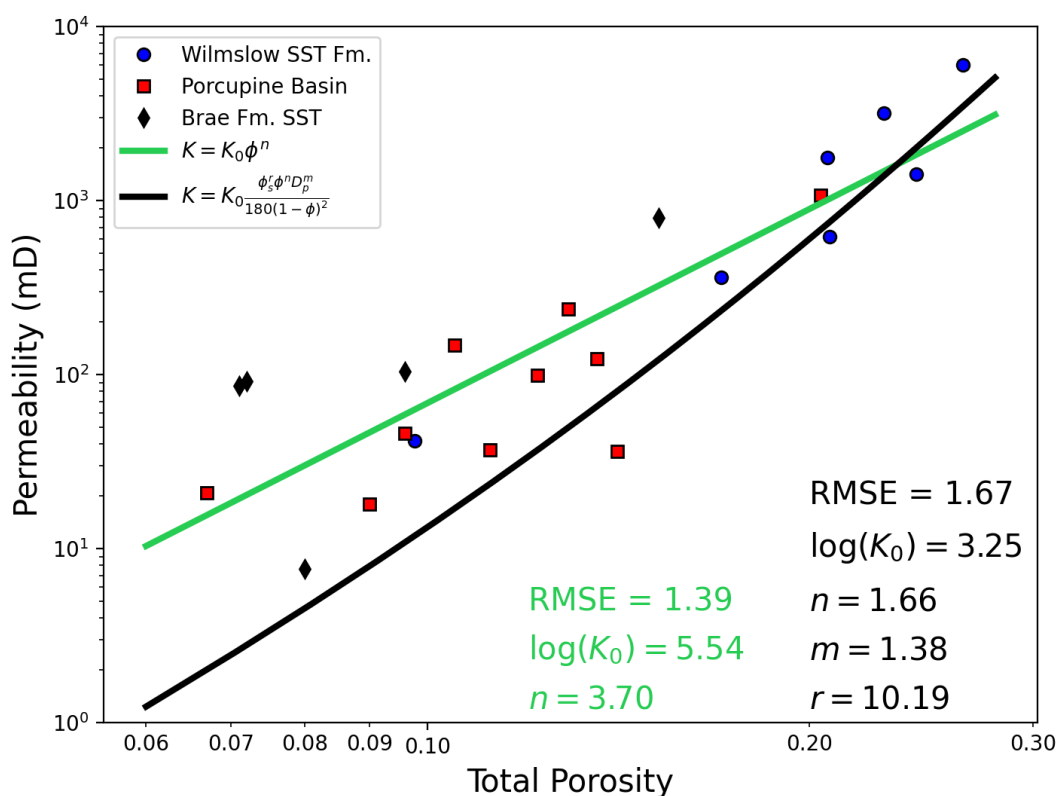
284 **Figure 5**

285 The best of the four fits based on the RMSE (Fig. 5) is the case where each
286 exponent can vary and is not constrained in any way, shown by the black fit line.
287 The red fit line, which omits the grain size term, produces the poorest quality fit
288 even though we identified grain size to have no relationship with porosity or
289 permeability (see Figs. 3a and 3c). The remaining two fit lines in cyan and orange
290 offer fits with RMSE values just larger and therefore less successful than the black
291 fit. The cyan and orange fits offer varying constraint on the exponents of grain size
292 alone and grain size alongside grain sphericity respectively but importantly, both
293 include the grain size term. Inclusion of this term, whether its exponent may vary
294 or not, clearly allows the given fit to be of a greater quality than omitting it all
295 together.

296 It is apparent that even the best fit achieved, shown by the black line in
297 Figure 5, does not fit all data points effectively, especially below a total porosity of
298 ca. 15%. Consequently, we show an additional, simpler fit which does not consider
299 any grain characteristics in Figure 6 (green line) alongside the best fit identified in
300 Figure 5. Our results show that the simpler fit which considers porosity and

301 permeability alone is more effective, exhibiting a lower RMSE of 1.39 as opposed
 302 to 1.67 in the case of the fit incorporating the grain characteristics.

303 **Figure 6**



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DISCUSSION

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Grain Boundary Watershed Segmentation

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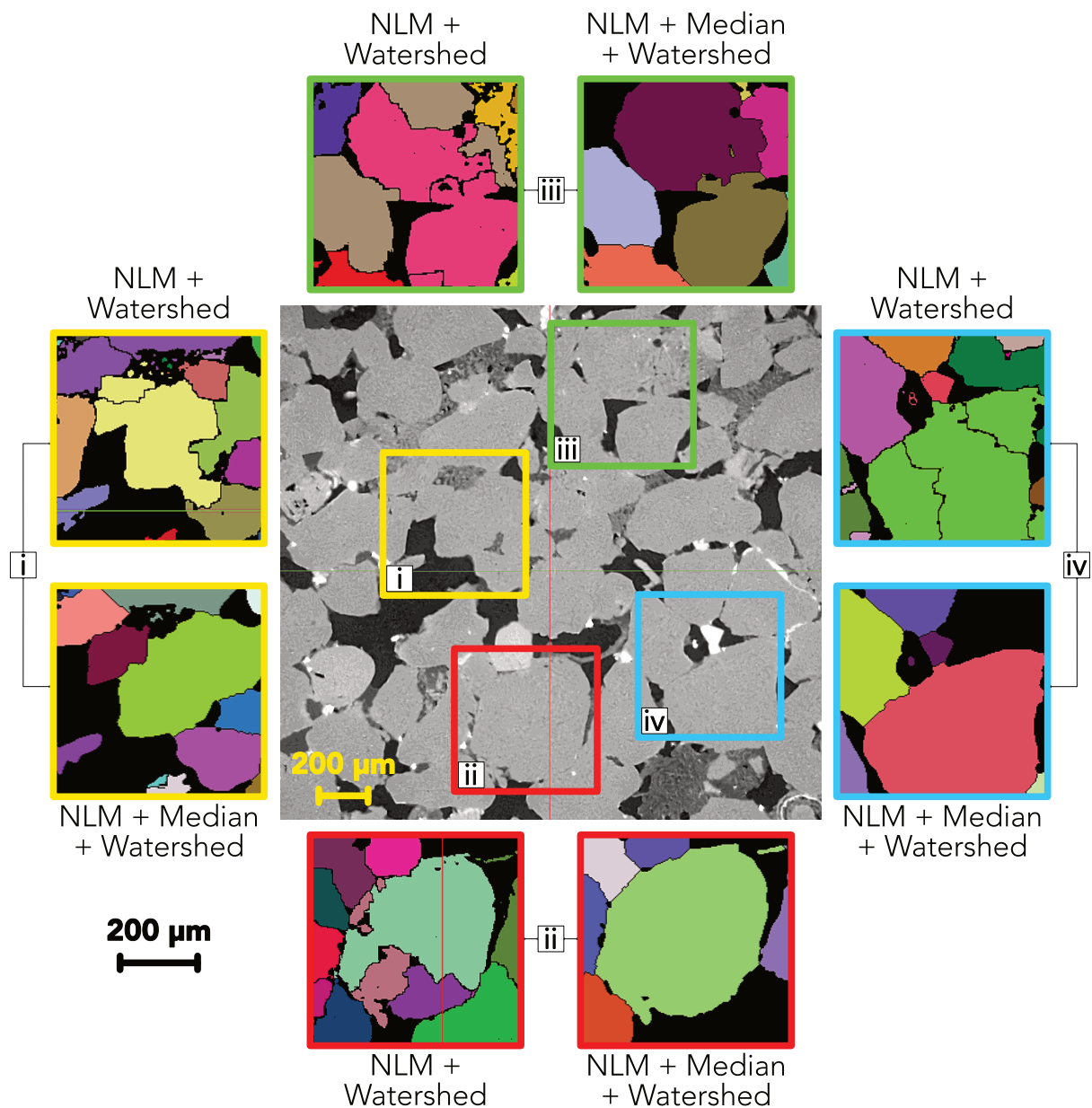
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Whilst the approach that we take to segment individual grains (Fei and Narsilio 2020) is relatively straightforward, we acknowledge that there are a number of other algorithms which aim to improve the accuracy of the traditional watershed algorithm. In particular, oversegmentation is an issue, particularly when segmenting features of a wide range of sizes and shapes where multiple

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312 markers are placed within one feature (Kong and Fonseca 2018; Sun et al. 2019;
313 Leonti et al. 2020). Modified watershed approaches have been developed using
314 the bring down method (Shi and Yan 2015) and the bring up method (Kong and
315 Fonseca 2018; Leonti et al. 2020) to accurately label features and their boundaries.
316 Due to the high accuracy of results reported by Fei & Narsilio (2020), the ease of
317 implementation and minimal computational cost we chose to use the traditional
318 watershed technique with a non-local means and median filter in line with the
319 methodology described by the authors.

320 The technique used here is very similar to that applied by Thomson, et al.
321 (2020a). Thomson, et al. (2020a) implement a traditional watershed algorithm but
322 only use a non-local means filter without a median filter. The non-local means
323 filter performs the bulk of the denoising in the images very effectively, but this
324 type of filter is not optimal for retaining or improving feature boundaries. In
325 contrast, the median filter is very effective for this purpose, enhancing the clarity
326 of feature boundaries whilst smoothing any remaining noise in the images. We
327 show the similarities and differences in the results of watershed segmentation
328 using the two approaches in Figure 7.



329 **Figure 7**

330 Our results show that the approach used by Thomson, et al. (2020a) results
331 in some oversegmentation of grains when comparing the watershed result to the
332 greyscale CT image. In contrast the approach used in this study does not show
333 severe oversegmentation of the same grains, owing to the boundary
334 enhancement provided by the median filter. Furthermore, by using the 3D Suite
335 plug-in for Fiji, grains which are touching the boundaries of the study volume can

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336 be excluded from measurement to ensure only grains which are complete and
337 truly representative are included. This was not included in the method used by
338 Thomson, et al. (2020a) and therefore partial grains may have significantly
339 influenced the mean grain measurements made.

340 Finally, Thomson, et al. (2020a) acknowledge in their work that the
341 separated grains in their work displayed an unexpected group of grains with Feret
342 diameters of $< 63 \mu\text{m}$, smaller than the classification of sand grains following the
343 scheme proposed by Wentworth (1922). Employing the additional median filter
344 largely removed the occurrence of these small, unexpected grains. Therefore, we
345 suggest that the combination of a median filter with a non-local means filter is
346 effective in reducing over segmentation and identification of small, unexpected
347 features.

348 *The Influence of Grain Characteristics*

349 **Grain Size and Shape.**--- The observed lack of relationships between mean
350 grain size and both porosity and permeability (Figs. 3a and 3c) strongly suggests
351 that grain size within this suite of samples is not influential on the porosity-
352 permeability relationship of the respective pore structures. Nabawy (2014)
353 presents a similar conclusion when examining the influence of grain size on
354 porosity and permeability in a series of idealised grain packs as well as in high
355 porosity sandstone samples. All but two of our samples are classified as very well-
356 well- or moderately-sorted (Folk 1980). Therefore, we suggest that future work

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357 should focus on the relationship between grain size and porosity and permeability
358 in a variety of sandstones of different grain maturity, shape and facies to identify
359 any factors which may influence whether grain size presents a relationship with
360 porosity or permeability.

361 In contrast, we show evidence that mean grain sphericity has a direct
362 positive impact on both porosity and permeability (Figs. 3b and 3d). Nabawy
363 (2014) identifies a similar relationship with the elongation (grain length/grain
364 diameter) of grains within their sample suite where less elongate grains
365 contribute to greater porosity and permeabilities. Nabawy (2014) uses elongation
366 as a measure of grain anisotropy where a more elongate grain indicates a greater
367 degree of anisotropy. We can apply the same approach to grain sphericity, where
368 a less spherical grain indicates a greater degree of anisotropy. Following this
369 paradigm, we see that our results agree with those of Nabawy (2014), a greater
370 degree of anisotropy of the grains results in a reduction in both porosity and
371 permeability.

372 We calculated a simple linear fit for the relationship between mean grain
373 sphericity and total porosity which is given by $\phi = 1.22 \phi_s - 0.42$. Nabawy (2014)
374 proposes a relationship between elongation, E and porosity using their sample
375 suite where $\phi = 45.73 E^{-1} + 9.19$. This provides two parameters by which a
376 porosity estimation may be made based upon two different measures of grain
377 anisotropy. Whilst Nabawy (2014) achieves an elongation fit exhibiting a

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378 correlation coefficient of 0.92 we find our sphericity fit to have a correlation
379 coefficient of 0.65. We consider three separate sample suites from different
380 sedimentary facies, whilst Nabawy (2014) focusses on a single sample suite, which
381 makes the relationship between anisotropy and porosity less clear. Consequently,
382 we suggest that different depositional environments may have a more significant
383 effect upon the characteristics which influence the relationship between grain
384 anisotropy and porosity, as opposed to there being one consistent relationship
385 being applicable across a wide variety of sandstones. Further research is required
386 to quantify the scale of this influence

387 We also investigated the control which the anisotropy of grains has on the
388 geometry of the pores themselves, finding that there is generally a positive
389 relationship between grain sphericity and pore diameter (Fig. 4). Our results agree
390 with the relationship identified between porosity and grain anisotropy, measured
391 through elongation (Nabawy 2014). This indicates that these two measures of
392 grain anisotropy exhibit similar controls on porosity which reflects directly in the
393 geometry of the pore structures.

394 A suggested limitation of the relationship reported by Nabawy (2014) is that
395 it may depend on grain elongation occurring systematically along one axis which
396 is common throughout the sampled material. Such imbrication of grains
397 according to their elongation axes may result due to the flow of depositional
398 currents and load pressure. Where such an alignment is not clearly present, for

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399 example under depositional conditions where turbulent flow dominates, these
400 results imply that the detrimental impact on permeability would be far more
401 pronounced than any influence on the relationship with porosity. This conclusion
402 requires further testing using samples from varied depositional environments to
403 eliminate the effects of sorting and stratification.

404 We observe an apparent group of seven outliers when examining the
405 relationship between grain sphericity and permeability (Fig. 3d) which fall below
406 the dominant trend. The fact that this group of outliers are not apparent when
407 comparing sphericity with porosity (Fig. 3b) strongly suggests that their rogue
408 placement is due to a factor which inhibits fluid flow but does not change the
409 absolute porosity measurement. This may point towards a lack of preferential
410 orientation with regards to grain anisotropy within these particular samples.

411 Further investigation of the seven outliers found that there was no
412 apparent common characteristic amongst the outliers which could differentiate
413 them from the remaining samples. We investigated whether there was a
414 relationship between these outliers and their sample depth, sorting, porosity or
415 permeability which might explain their occurrence. None of these characteristics
416 helped to explain the presence of the seven outliers. Furthermore, a qualitative
417 assessment of the μ CT images found nothing of significance which might allow for
418 the differentiation of this sample group such as presence of cement or other
419 precipitates which were not present in the main group of samples.

420 It might be expected that a lack of grain orientation would manifest
421 throughout a given geological unit, leading to surprise that the outlier group
422 contains at least one sample from each of the three studied formations. We
423 suggest that the resulting texture may be controlled by a different depositional
424 process. Alternatively, the scale of the sample upon which measurements were
425 made could be considered not suitably representative for the scale of the
426 processes which cause variation in grain imbrication and alignment with regards
427 to anisotropy. Therefore, we suggest that future work should focus on identifying
428 a suitable representative elementary volume over which measures of grain
429 anisotropy, such as elongation and sphericity, can be representatively measured.
430 Equally, identification and implementation of a technique to measure and
431 quantify alignment or imbrication of grains in 3D at the pore scale would be
432 beneficial in providing greater context for relationships between porosity and
433 permeability with measures of grain anisotropy.

434 **Grain Influence on the Porosity-Permeability Relationship.**--- Despite
435 the positive relationship identified between mean grain sphericity and porosity
436 and permeability (Figs. 3b and d) we have found that the influence of grain
437 characteristics is not beneficial to constraining the porosity-permeability
438 relationship in these sample suites (Fig. 6). This may be a result of using a Kozeny-
439 Carman fit equation which makes the assumption that grains are spherical
440 producing a simple pore structure (Rahrah et al. 2020). Bear (1972) describes how

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441 this assumption arises from the transformation of the specific surface area term
442 (Carman 1937) to a characteristic grain size term.

443 Inclusion of grain size in the paradigm of a Kozeny-Carman relationship
444 defines the diameter of the grain which is assumed to be spherical. However, we
445 define grain size as the greatest distance from one side of the grain to another or
446 the calliper distance, which is applicable to non-spherical grains. Therefore, as the
447 sphericity of a given grain reduces, it moves further from the Kozeny-Carman
448 assumption which results in a poorer fit to samples with a lower mean grain
449 sphericity. We show that a lower sphericity results in a lower porosity and
450 permeability (Fig. 3) therefore, we would expect the Kozeny-Carman fit to be
451 poorer at lower porosities and permeabilities.

452 We show it to be the case that lower sphericity or greater grain anisotropy
453 results in a poorer agreement with a Kozeny-Carman based fit (Figs. 5 and 6). It
454 can be observed that below ca. 15% total porosity just one data point lies below
455 the fit line whereas the remaining data points lie consistently and significantly
456 above the fit lines calculated using equation 6. For example, sample PB12 has a
457 low mean grain sphericity of 0.37 and a relatively low total porosity of 9% and can
458 be seen to plot above the black K-C fit line (Fig. 6). This strongly suggests that the
459 Kozeny-Carman style fit is not suitable for use with samples which possess grains
460 which show significantly low sphericities. Torskaya et al. (2014) investigate the
461 effect of grain shape on permeability and find that when using realistic grain

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462 shapes from μ CT images that the K-C equation underestimates permeability by
463 between 30 and 70%. When using simplified and spherical grain shapes Torskaya
464 et al. (2014) find that the K-C equation fit was far more successful, supporting our
465 conclusion that the K-C spherical grain assumption is causing the poor quality fit.
466 The K-C approach therefore, is not suitable for use with materials where grains
467 are significantly non-spherical.

468 As a result of this identified limitation, we propose that future work should
469 look to develop an alternative model which accounts for variation in grain
470 sphericity within and between different sandstone samples. In this study we have
471 clearly shown that grain sphericity exhibits a strong relationship with both
472 porosity and permeability (Fig. 3), highlighting the possible value in incorporating
473 this grain characteristic in a porosity-permeability model. A model which is still
474 able to incorporate each influencing factor as individual terms (as in equation 6)
475 would be favourable to provide flexibility and the ability for experimentation. Such
476 a model could be tested against the simple and K-C models presented in Figure 6
477 based upon RMSE.

478 Whilst many modified versions of the Kozeny-Carman equation have been
479 proposed and used (e.g., Le Gallo et al. 1998; MacQuarrie and Mayer 2005;
480 Hommel et al. 2018), the fundamental assumption of spherical grains and pores
481 arranged as bundles of capillaries remains. Alternatives to a K-C approach at the
482 same scale have been used to describe permeability such as the Fair-Hatch,

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483 Brinkman and Panda and Lake models, described and summarised by Le Gallo et
484 al. (1998) and MacQuarrie & Mayer (2005). Whilst some of these approaches use
485 grain size terms, they do not include terms which allow for direct inclusion of grain
486 shape or anisotropy.

487 A further consideration which would be highly beneficial to any future
488 model would be to account for the percolation threshold, a key phenomenon
489 which makes effectively characterising the porosity-permeability relationship
490 difficult over a range of porosities. Thomson, et al. (2020b) and Payton et al. (2021)
491 show the percolation threshold for full connectivity to be at ca. 8 - 15% total
492 porosity, whilst Mavko & Nur (1997) and Rahrah et al. (2020) show the value of
493 incorporating the percolation threshold into a K-C style fit. Consideration of the
494 percolation threshold alongside variable grain sphericity would surely be an
495 effective approach to best describe the porosity-permeability relationship.

496

497

CONCLUSIONS

498 In this work we made a comparison of two similar grain segmentation
499 techniques, using marker-based watershed algorithms, for reliable and accurate
500 grain boundary identification across our sample suites. We found that using a
501 median filter in addition to a non-local means (NLM) filter prior to segmentation
502 resulted in superior grain separation as opposed to using a NLM filter alone. This
503 appeared to be due to the ability of the median filter to preserve and enhance the

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504 grain edges during denoising, reducing oversegmentation. The low computational
505 cost and high speed at which this technique can be applied makes this a suitable
506 option for segmentation of sandstone materials such as those investigated here.

507 We have used digital image analysis techniques on μ CT images of three
508 different suites of sandstone samples to investigate the impact of grain
509 characteristics on the porosity-permeability relationship. We have shown that in
510 this collection of samples the porosity-permeability relationship is not better
511 constrained when including grain shape and size parameters in a Kozeny-Carman
512 type fit equation. This is the case despite identification of a strong positive
513 relationship between grain sphericity and both porosity and permeability. We
514 found no such relationship with grain size. Therefore, we found a porosity-
515 permeability relationship best described by $K = 10^{5.54} \phi^{3.7}$.

516 We determine that the need to assume that grains are spherical when
517 working in a Kozeny-Carman paradigm is severely limiting to identifying an
518 effective porosity-permeability relationship. Future work should focus on
519 incorporating a grain sphericity term in a model which effectively handles non-
520 spherical and non-uniform grains. Of added benefit would be consideration of the
521 percolation threshold in producing a model capable of constraining the porosity-
522 permeability relationship over a range of porosities in sandstones.

523 Finally, consideration of grain sphericity as a measure of 3D grain shape
524 anisotropy revealed a relationship of decreasing anisotropy resulting in greater

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525 porosity and permeability, in agreement with 2D measures of grain anisotropy.
526 We found total porosity to vary with grain sphericity according to $\phi = 1.22\phi - 0.42$,
527 offering an additional indirect method of predicting porosity. A group of outliers
528 are identified, vertically displaced below the main trend of the sphericity-
529 permeability data. We suggest that this may be due to a lack of grain orientation
530 with regards to sphericity in these samples, inhibiting the permeability only as the
531 same occurrence is not observed so strongly in the case of porosity.

532

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668 **CONFLICT OF INTEREST STATEMENT**

669 The authors declare that the research was conducted in the absence of any
670 commercial or financial relationships that could be construed as a potential
671 conflict of interest.

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AUTHOR CONTRIBUTIONS

674

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675

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692

DATA AVAILABILITY STATEMENT

693 The μ CT images used in this article are available from a variety of sources.
694 Images of the Wilmslow Sandstone Fm. for samples with a SF prefix are available
695 from Payton et al. (2021). Images of the Brae Fm. Sandstone for samples with BFS
696 prefix are not publicly available and must be requested from Thomson, et al.
697 (2020b). Images of the Minard Formation Sandstone from the Porcupine Basin for
698 samples with a PB prefix are available from the Royal Holloway, University of
699 London Figshare Repository, <https://figshare.com/s/a3c53f2b89fb3d655f6a>.

700

701

FIGURE CAPTIONS

702

Figure 1. Schematic diagram showing the typical steps in grain

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identification using a watershed technique on CT images.

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Figure 2. Isolated collection of grains (white) and single grain (orange)

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shown in 3D from sample SF696. The saw-tooth or staircase pattern is

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highlighted which arises from the voxelised images. This can lead to

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overestimation of surface area and impact the subsequent sphericity

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measurements.

709

Figure 3. Relationship between mean grain size, mean grain sphericity

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and total porosity and permeability. A generally positive relationship with

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porosity and permeability can be observed in the case of mean grain sphericity

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but no such relationship is present with mean grain size. A region of outliers is

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identified by a dashed line in **(d)** with the same data points also identified in **(b)**.

714 **Figure 4.** Relationship between mean grain sphericity and mean
715 connected pore diameter for each of the three sample suites. It is apparent that
716 there is a generally positive relationship between the two parameters.

717 **Figure 5.** Range of calculated fit configurations to the porosity-
718 permeability relationship which incorporate grain characteristics using a Kozeny-
719 Carman based relationship. The table to the right qualitatively describes the
720 difference between each fit line whilst the respective equations are displayed in
721 the plot legend.

722 **Figure 6.** Calculated fits to the porosity-permeability relationship. The root
723 mean square error (RMSE) values are reported for each fit, showing that the
724 better fit is the simpler one in green. The green fit excludes any measured grain
725 characteristics whereas the black fit does not.

726 **Figure 7.** Comparison of two different filtering techniques' effect on the
727 watershed algorithm in a single slice of sample PB10. Four different locations
728 have been highlighted for comparison on an image which has undergone non-
729 local means (NLM) filtering only. Annotated squares show the result of
730 watershed grain segmentation following only NLM (Thomson et al. 2020a) and
731 NLM with a median filter (Fei and Narsilio 2020). Each grain can be identified by
732 a different colour however, due to the number of grains, colours have been
733 reused and instead the black grain boundaries split different grains of the
734 same colour. In each annotation an example of over-segmentation is observed

735 in the case of using NLM filtering only when compared to what we might expect
 736 from the CT image. The outer scale bar applies to all annotations.

737

738 **Table 1.** Summary of the sampled materials analysed in this study.

739

740 *Payton et al. (2021), **Thomson et al. (2020b).

741

Sampling Location	Well ID	Sample ID	Depth (m)	Geology	
Porcupine Basin, N. Atlantic	26/28-1	PB01 PB02 PB03 PB05	2271 2256.4 2420 2420.48	Minard Formation	Renard Member
	26/28-2	PB06 PB07 PB08 PB10 PB11 PB12	2117 2118 2116.8 2118.6 2119.15 2119.85		Dooneragh Member
Sellafield, UK*	SFBH13 B	SF696 SF697 SF698 SF699 SF700 SF701 SF702	63.8 76.1 96.98 126.27 144.03 172.16 181.39	Wilmslow Sandstone Formation	
North Sea, UK**	16/7b-20	BFS1 BFS2 BFS4	4040.1 4041.35 4045.13	Brae Formation Sandstone	
	16/7b-23	BFS5 BFS8	4061 4063.75		

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744 **Table 2.** Grain-based measurements made for each sample.

745

Sample	Sorting (ϕ)	Mean Grain Size (μm)	Mean Grain Sphericity
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PB01	0.63	242	0.45
PB02	0.61	298	0.43
PB03	0.44	112	0.47
PB05	0.45	297	0.47
PB06	0.55	198	0.46
PB07	0.45	92	0.44
PB08	0.42	168	0.48
PB10	0.49	120	0.45
PB11	0.78	223	0.40
PB12	0.56	117	0.37
SF696*	0.61	203	0.49
SF697*	0.54	205	0.46
SF698*	0.64	204	0.52
SF699*	0.50	257	0.53
SF700*	0.51	230	0.46
SF701*	0.51	179	0.50
SF702*	0.52	247	0.45
BFS1**	0.61	135	0.44
BFS2**	0.75	262	0.43
BFS4**	0.69	158	0.44
BFS5**	0.53	421	0.43
BFS8**	0.44	108	0.46

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747 *Payton et al. (2021), **Thomson et al. (2020b).

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757 **Table 3.** Porosity and permeability measurements made for each sample.
758

Sample	Total Porosity (%)	Connected Porosity (%)	Permeability (mD)
PB01	20.4	20.3	1070
PB02	10.5	9.8	147
PB03	12.2	10.2	99
PB05	11.2	4.9	37
PB06	6.7	5.2	21
PB07	9.6	8.9	46
PB08	13.6	13.3	123
PB10	12.9	9.7	237
PB11	14.1	13.6	36
PB12	9	6.9	18
SF696*	20.7	20.4	1760
SF697*	20.7	20.3	620
SF698*	22.9	22.7	3190
SF699*	26.4	26.3	6040
SF700*	17.0	16.6	360
SF701*	24.3	24.1	1420
SF702*	9.77	8.89	40
BFS1**	7.2	5.8	91
BFS2**	7.1	5.7	86
BFS4**	9.6	9.1	104
BFS5**	7.8	5.1	6.7
BFS8**	15.2	14.8	795

759
760 *Payton et al. (2021), **Thomson et al. (2020b).