1	Estimating 'outbreak'-level tornado counts and casualties from
2	environmental variables
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## ABSTRACT

Environmental variables are routinely used in estimating when and where tornadoes are likely to 9 occur, but more work is needed to understand how tornado and casualty counts of severe weather 10 outbreak vary with the larger scale environmental factors. Here the authors demonstrate a method 11 to quantify 'outbreak'-level tornado and casualty counts with respect to variations in large-scale 12 environmental factors. They do this by fitting negative binomial regression models to cluster-level 13 environmental data to estimate the number of tornadoes and the number of casualties on days with 14 at least ten tornadoes. Results show that a 1000 J kg<sup>-1</sup> increase in CAPE corresponds to a 5% 15 increase in the number of tornadoes and a 28% increase in the number of casualties, conditional 16 on at least ten tornadoes, and holding the other variables constant. Further, results show that 17 a 10 m s<sup>-1</sup> increase in deep-layer bulk shear corresponds to a 13% increase in tornadoes and a 18 98% increase in casualties, conditional on at least ten tornadoes, and holding the other variables 19 constant. The casualty-count model quantifies the decline in the number of casualties per year and 20 indicates that outbreaks have a larger impact in the Southeast than elsewhere after controlling for 21 population and geographic area. 22

# 23 1. Introduction

Estimating tornado and casualty counts (hereafter referred to as characteristics) of severe weather 24 outbreaks is an important and challenging problem. It is important because of the potential for 25 loss of life and property damage. It is challenging because of the uncertainties associated with 26 exactly how many and where the tornadoes will occur. But progress is being made. Guidance from 27 dynamical models help forecasters outline areas of possible severe weather threats days in advance 28 (Hitchens and Brooks 2014) while guidance from statistical models help forecasters quantify 29 probabilities for given severe weather events (Thompson et al. 2017; Cohen et al. 2018; Elsner 30 and Schroder 2019; Hill et al. 2020). For example, Cohen et al. (2018) use a regression model 31 to specify the probability of tornado occurrence given certain environmental and storm-scale 32 conditions (circulation above radar level, rotational velocity, circulation diameter, etc). Elsner 33 and Schroder (2019) extend this model by making use of the cumulative logistic link function 34 that estimates probabilities for each damage rating using storm-relative helicity, bulk shear, and 35 convective available potential energy (CAPE). These studies put statistical guidance for estimating 36 severe weather outbreak characteristics on a firm mathematical foundation (Cohen et al. 2018; 37 Elsner and Schroder 2019). Room for additional work in this area motivates the present study. For 38 instance, the cumulative logistic regression (Elsner and Schroder 2019) provides a distribution for 39 the *percentage* of tornadoes within each Enhanced Fujita (EF) rating category (Fujita 1981), but a 40 model is needed to estimate the overall number of tornadoes given the likelihood of at least some 41 tornadoes. 42

Tornado outbreaks pose a risk of significant loss of life and property. Anderson-Frey and Brooks (2019) consider the role environmental factors play in the number of outbreak fatalities. They use self-organizing maps on the significant tornado parameter (STP) and find that more damaging

tornadoes (>EF3) present a higher risk for fatalities. However, they also note that both deadly and 46 non-deadly tornadoes are associated with high values of STP. Self-organizing maps are useful for 47 describing the role of environmental variables on casualties, but a statistical model is needed to 48 quantify the relationship between casualty counts and environmental factors. Here we demonstrate 49 a method to model 'outbreak'-level tornado and casualty counts from environmental conditions 50 and predefined tornado clusters. A cluster is defined (informally) as a group of ten or more 51 tornadoes occurring over a relatively short time scale (e.g., one day) and over a relatively limited 52 spatial domain (e.g., one to three states) (Fig. 1). The model allows us to quantify the associative 53 relationships between environmental variables and tornado counts. Moreover, the approach might 54 eventually help extend the available statistical guidance for predicting outbreak characteristics 55 particularly when combined with other models. 56

In this paper, we focus on outbreaks rather than on individual tornadoes. The larger space and time scales associated with the outbreak matches our interest in the larger-scale environmental factors like CAPE and shear. In what follows, we call the outbreaks 'clusters' as is done in Schroder and Elsner (2019) because we make no attempt to associate the cluster with a particular weather system. We do not reference the weather feature(s) responsible for the cluster although it is likely that many of them are associated with a single synoptic-scale system.

<sup>65</sup> Clusters in the United States are most frequent during April, May, and June (Dixon et al. 2014; <sup>64</sup> Tippett et al. 2012; Dean 2010) with most of them occurring across the Central Plains and the <sup>65</sup> Southeast. Clusters are less common in the Southeast and the Southern Plains during the summer <sup>66</sup> months as the jet stream migrates north taking the necessary wind shear with it (Concannon et al. <sup>67</sup> 2000; Gensini and Ashley 2011; Jackson and Brown 2009). The percentage of all tornadoes <sup>68</sup> occurring in clusters has recently been found to be increasing over time (Moore 2017; Tippett et al. <sup>69</sup> 2016; Elsner et al. 2015; Brooks et al. 2014).

This paper has two objectives: (1) demonstrate that environmental conditions prior to the 70 occurrence of any tornadoes can be used to skillfully model the number of tornadoes in a cluster 71 containing at least ten tornadoes (tornado-count model), and (2) show that these same environmental 72 conditions can be used to estimate the number of casualties if the number of people in harm's 73 way is known (casualty-count model). We accomplish these objects by fitting negative binomial 74 regressions to cluster-level data including environmental variables and tornado and casualty counts 75 on a convective day (24-hour period between 12 UTC and 12 UTC) when the number of tornadoes 76 is at least ten [see Elsner and Schroder (2019)]. The paper is outlined as follows. The data and 77 methods are discussed in section 2 including the mathematics of a negative binomial regression. 78 Statistics describing the response (i.e., tornado-casualty counts) and environmental variables are 79 given in section 3. The modeling results are presented in section 4, and a summary with conclusions 80 are given in section 5. 81

## **2.** Data and methods

We fit regression models to a set of tornado and reanalysis data aggregated to the level of tornado 83 clusters. Here, we describe how we organize the data and the procedures to aggregate values to the 84 cluster level. For our purposes, a cluster is a group of at least ten tornadoes occurring relatively 85 close to one another in both space and time between over a convective day. Ten is chosen as a 86 compromise between too few clusters leading to greater uncertainty and too many clusters leading 87 to excessive time required to fit the models (Elsner and Schroder 2019). Ten is also the number 88 that is sometimes used formally to define an outbreak (Galway 1977; Anderson-Frey et al. 2018). 89 The number of tornadoes in each cluster is the response variable in the tornado-count regression 90 model, and the number of casualties is the response variable in the casualty-count regression model. 91

Explanatory variables include outbreak size and location as well as environmental variables from
 reanalysis data representing conditions before the occurrence of the first tornado in the cluster.

### 94 a. Tornado clusters

First, we extract the date, time, genesis location, and magnitude of all tornado reports between 1994 and 2018 from the Storm Prediction Center [SPC] (https://www.spc.noaa.gov/gis/ svrgis/). We choose 1994 as the start year because it is the first year of the extensive use of the WSR-88D Radar (Heiss et al. 1990). In total, there are 30,497 national tornado reports during this period. The geographic coordinates for each genesis location are converted to Lambert conic conformal coordinates, where the projection is centered on 96° W longitude.

Next, we assign to each tornado a cluster identification number based on the space and time 101 differences between genesis locations. Two tornadoes are assigned the same cluster identification 102 number if they occur close together in space and time (e.g., 1 km and 1 h). When the difference 103 between individual tornadoes and existing clusters surpasses 50,000 s (~ 14 h), the clustering ends. 104 The space-time differences have units of seconds because we divide the spatial distance by  $15 \text{ m s}^{-1}$ 105 to account for the average speed of tornado-producing storms. This speed is commensurate with 106 the magnitude of the steering-level wind field across the clusters. The clustering is identical to 107 that used in Elsner and Schroder (2019) who developed a cumulative logistic model to the damage 108 scale at the individual tornado level. Additional details on the procedure, as well as a comparison 109 of the identified clusters to well-known outbreaks, are available in Schroder and Elsner (2019). 110

We keep only clusters having at least ten tornadoes occurring within the same convective day, which results in 768 clusters with a total of 17,069 tornadoes. The average number of tornadoes per cluster is 22 and the maximum is 173 (April 27, 2011). There are 80 clusters with exactly ten tornadoes. Each cluster varies by area and by where it occurs geographically (see Fig. 1 for examples of clusters). The cluster area is defined by the minimum convex hull (black polygon) that includes all the tornado genesis locations. The July 19, 1994 cluster with nine tornadoes over northern Iowa and one over northwest Wisconsin had an area of 33,359 km<sup>2</sup> and lasted about four hours. The April 27, 2011 cluster had 173 tornadoes spread over more than a dozen states and had an area of 1,064,337 km<sup>2</sup> with tornadoes occurring throughout the 24-h period (12-UTC to 12-UTC).

For each cluster, we sum the number of injuries and deaths across all tornadoes to get the clusterlevel number of casualties (sum of injuries and fatalities). Further, we estimate the population within the cluster area and the geographic center of the cluster. Population values are U.S. Census Bureau estimates in cities with at least 40,000 people (Steiner 2019). Population is used as an explanatory variable in place of cluster area in the casualty-count model.

We encountered situations where there were multiple clusters on a given day. For example, there were two clusters on November 11, 1995 (Fig. 2). The first cluster was responsible for 15 tornadoes resulting in three casualties. The second cluster was responsible for 10 tornadoes resulting in four casualties. Both clusters were the result of a single cold front along the eastern coast of the United States. However, they are not grouped into a single cluster because the minimum between-tornado space-time distance was larger than our threshold.

## <sup>132</sup> b. Environmental variables

Large-scale environmental conditions for producing tornadoes are well studied and include large magnitudes of convective available potential energy, bulk shear, and weak convective inhibition (Brooks et al. 1994; Rasmussen and Blanchard 1998; Thompson et al. 2003; Shafer and Doswell 2011; Doswell et al. 2006). We obtain variables associated with these environmental conditions from the National Centers for Atmospheric Research's North American Regional Reanalysis <sup>138</sup> (NARR), which is supported by the National Centers for Environmental Prediction (Mesinger et al. <sup>139</sup> 2006). Each variable has numeric values given on a 32-km raster grid with the values available <sup>140</sup> in three-hour increments starting at 00 UTC. In the severe weather literature, these environmen-<sup>141</sup> tal variables are called 'parameters'. However here, since we employ statistical models, we call <sup>142</sup> them variables to be consistent with the statistical literature where the word 'parameter' denotes <sup>143</sup> unknown model coefficients and moments of statistical distributions (e.g., the mean).

We select environmental variables at the nearest three-hour NARR time *prior* to the occurrence 144 of the first tornado in the cluster. For example, we use the environmental variables given at 15 UTC 145 if the first tornado in a cluster occurs at 16:30 UTC. This selection criteria results in a sample of the 146 environment that is less contaminated by the deep convection itself but at a cost that underestimates 147 the severity in cases where environmental conditions rapidly change favoring tornado development. 148 About 60% of all clusters have the initial tornado occurring between 18 and 00 UTC (Table 1). 149 However, there are more tornadoes in clusters when the first tornado occurs between 15 and 18 150 UTC on average. 151

The environmental variables we consider include convective available potential energy (CAPE) 152 and convective inhibition (CIN) as computed using the near-surface layer (0 to 180 mb above the 153 ground level) consistent with Allen et al. (2015b). We also include deep (1000 to 500 mb) and 154 shallow (1000 to 850 mb) layer bulk shears (DLBS, SLBS) computed as the square root of the sum 155 of the squared differences between the u and v wind components at the respective levels consistent 156 with Tippett et al. (2012). Climate researchers use these NARR variables at the climatological 157 scale as proxies for the more traditional variables used in forecasting severe weather (Allen et al. 158 2015b; Moore et al. 2016; Tippett et al. 2012). 159

We take the highest (lowest for CIN) value across the grid of values within the area defined by the cluster's convex hull (Fig. 3). This is done to capture environmental conditions that represent the <sup>162</sup> unadulterated pre-tornado environment. In contrast, the mean (or median) value is influenced by <sup>163</sup> conditions throughout the domain including earlier occurring non-tornado-producing convection <sup>164</sup> and in areas within the clusters that did not experience tornado activity. Histograms of the <sup>165</sup> maximums (not shown) show no evidence of extreme behavior.

We do not include storm-relative helicity (SRH), lifted condensation level (LCL), or dewpoint 166 temperature (DEW) in this research although these variables have proven to be indicators of 167 favorable environments for tornado production. Storm-relative helicity is not used because it is 168 correlated with DLBS and SLBS (Table 2). Likewise dew-point temperature and LCL height 169 are not used because of their relatively high correlation with CAPE. A model that included SRH, 170 DEW, and LCL height as predictors showed no significant improvement over a model without them. 171 Further, we do not use composite variables including the significant tornado parameter (STP) and 172 the supercell composite parameter (SCP). STP, for example, is the product of variables including 173 CAPE, storm-relative helicity, CIN, and lifted condensation level (LCL) height. A moderate value 174 of STP can result from either high CAPE and low shear or low CAPE and high shear environments 175 holding the other variables constant. Here we separate this composite relationship to examine the 176 direct relationships between CAPE and shear on tornado activity at the scale of outbreaks. 177

## <sup>178</sup> c. Negative binomial regression

We fit a negative binomial regression model to the cluster-level data. We chose this type of regression because the response variable in the tornado (and casualty) model is a count. A count variable is described by a discrete distribution like the Poisson or negative binomial rather than by a continuous distribution like the Normal (Gausssian). The choice of which discrete distribution is made in favor of the negative binomial since the mean number of tornadoes (casualties) per cluster is substantially smaller than the variance in the number of tornadoes (casualties), whereas
 the Poisson distribution assumes the mean is equal to the variance.

With the cluster as our unit of analysis, we fit a series of negative binomial regression models to the data having the form

$$T \sim \text{NegBin}(\hat{\mu}, n)$$

$$\ln(\hat{\mu}) = \beta_0 + \beta_A A + \beta_\phi \phi + \beta_\lambda \lambda + \beta_Y Y +$$

$$\beta_{CAPE} \text{CAPE} + \beta_{CIN} \text{CIN} + \beta_{DLBS} \text{DLBS} + \beta_{SLBS} \text{SLBS},$$
(1)

where the number of tornadoes (T) is the dependent variable that is assumed to be adequately described by a negative binomial distribution (NegBin) with a rate parameter  $\mu$  and a size parameter 189 *n* (Hilbe 2011). The natural logarithm of the rate parameter is linearly related to cluster area (A), 190 cluster center location [latitude ( $\phi$ ) and longitude ( $\lambda$ )], year (Y) and the four environmental variables 191 (CAPE, CIN, DLBS, and SLBS). These are the explanatory variables. The model is fit using the 192 method of maximum likelihoods carried out in the call to the glm.nb function from MASS package 193 in R (Venables and Ripley 2002). We do the same for the initial casualty-count model, but we 194 replace cluster area with population (P). We simplify the initial models through single-term 195 deletions as described in §4. 196

Regression model skill is evaluated by comparing the observed tornado and casualty counts with what is predicted by the model. The predicted rates for each cluster are obtained by plugging the values of the associated explanatory variables into the model. Predicted rates are under dispersed (lower variation) relative to the observed counts. Comparisons are made using the metrics of Pearson correlation coefficient and mean absolute error. Predictive skill using these metrics is evaluated using in-sample and out-of-sample predictions. In-sample predictions are made using all clusters to fit a single model while out-of-sample predictions are made by successively holding one cluster out of the model fitting procedure and using the particular model to predict the counts

<sup>205</sup> from the cluster left out [hold-one-out cross validation; see Elsner and Schmertmann (1994)].

# 206 **3. Results**

## 207 a. Descriptive statistics

The number of clusters decreases exponentially with an increasing number of tornadoes per 208 cluster (Fig. 4). There are 80 clusters with ten tornadoes but only ten clusters with 30 tornadoes. 209 The right tail of the count distribution is long with the April 27, 2011 cluster having 173 tornadoes 210 [47 (6%) of the clusters have more than 50 tornadoes and are not shown]. However, more clusters 211 have 20 or 21 tornadoes than expected from a simple decay function. This deviation is unlikely 212 the result of physical processes, and it appears too large to be sampling variability. It might be 213 due to a consistent rounding of the totals to the nearest five or ten. There is an upward trend in 214 the number of tornadoes per cluster (not shown) consistent with recent studies (Elsner et al. 2015). 215 The distribution of casualties is also skewed toward many clusters having only a few casualties and 216 a few have many. Thirty-six percent of all clusters (275) are without a casualty and 56% of the 217 clusters have fewer than four casualties. 218

There is a seasonality to the chance of at least one tornado cluster (Fig. 5). The empirical seven-day probability of at least one cluster is between 20 and 30% for much of the year except between the middle of March and early July (Fig. 5A). The probabilities approach 80% between mid and late May. The number of tornadoes per cluster is less variable ranging between about 10 and 35 tornadoes per week with no strong seasonality although clusters during July and August tend to have somewhat fewer tornadoes (Fig. 5B). The casualty rate, defined as the number of casualties per 100,000 people within the cluster area, has a distinct seasonality with rates being
highest between March to April and August to September (Fig. 5C).

Across the 768 clusters the mean of the maximum values of CAPE is 2,225 J kg<sup>-1</sup> and the mean of the minimum values of CIN is -114 J kg<sup>-1</sup> (Table 3). The maximum deep-layer bulk shear values range from 5.6 to 47.9 m s<sup>-1</sup>. Cluster areas range from 361 to 1,064,337 km<sup>2</sup> with an average of 167,990 km<sup>2</sup>.

#### <sup>231</sup> b. A model for the number of tornadoes

First, we fit a negative binomial regression to the cluster-level tornado counts using the explanatory variables given in Table 3. This is our tornado-count model. We divide the cluster area by 10 million so it has units of 100 km<sup>2</sup>. We divide CAPE by 1000 so it has units of 1000 J kg<sup>-1</sup> and we divide CIN by 100 so it has units of 100 J kg<sup>-1</sup>. This simplifies interpretation of the model coefficients, but does not affect the goodness of fit.

All terms have signs on the coefficient that are physically reasonable (Table 4). The number of 237 tornadoes in a cluster increases with cluster area, CAPE, and bulk shear (deep and shallow layers) 238 and increases for decreasing CIN (i.e., less inhibition) as expected. The significance of the variable 239 in statistically explaining tornado counts is assessed by the corresponding z-value given as the ratio 240 of the coefficient estimate to its standard error (S.E.). We reject the null hypothesis that a particular 241 variable has no explanatory power if its corresponding *p*-value is less than .01. Here we fail to 242 reject the null hypothesis for the variables latitude, longitude, and year, which indicates that these 243 non-physical variables have a relatively small impact on tornado counts relative to the physical 244 variables given the data and the model. In particular, there is no significant trend over time in the 245 number of tornadoes in these clusters. The only physical variable that is not statistically significant <sup>247</sup> is CIN. This is likely a result of the NARR data being too coarse to adequately represent CIN. We
 <sup>248</sup> remove all statistically insignificant variables before fitting a final model.

All variables in the final model are significant although the magnitudes of the coefficients have 249 changed a bit relative to their values in the initial model. The in-sample correlation between the 250 observed counts and predicted rates is .59 [(0.54, 0.64), 95% uncertainty interval (UI)] (Fig. 6). 251 We find that the model is not improved by using the average values of these same environmental 252 variables. The model statistically explains almost 60% of the variation in cluster-level tornado 253 counts but tends to over predict the number of tornadoes for smaller clusters and slightly under 254 predict the number of tornadoes for larger clusters. We test the significance of area and find the 255 model performance decreases by 25% when excluding area from the model. The mean absolute 256 error between the observed counts and expected rates is 8.6 tornadoes or 5.2% of the range in 257 observed counts and 9.3% of the range in predicted rates. The out-of-sample errors are quite 258 similar due to the large sample size (768 clusters). A hold-one-out cross validation exercise (Elsner 259 and Schmertmann 1994) results in an out-of-sample correlation of .58 and a mean absolute error 260 of 8.6 tornadoes. The lag-1 temporal autocorrelation in cluster-level tornado counts is .13. 261

The value of  $\beta_0$  (Table 4) is the regression estimate when all variables in the model are evaluated at 262 zero. The effect size for a given explanatory variable is given by the magnitude of its corresponding 263 coefficient. The coefficient is expressed as the difference in the logarithm of the expected tornado 264 counts for a unit increase in the explanatory variable holding the other variables constant. For 265 example, the scaled units of CAPE are 1000 J kg<sup>-1</sup>. An increase in CAPE of 1000 J kg<sup>-1</sup> results in 266 a  $(\exp(.0459) - 1) \times 100\% = 4.7\%$  increase in the expected number of tornadoes, conditional on at 267 least ten tornadoes. Continuing, units of deep-layer bulk shear are 10 m s<sup>-1</sup> so an increase in shear 268 of 10 m s<sup>-1</sup> results in a 13% increase in the expected number of tornadoes. A similar increase in 269 shallow-layer bulk shear results in a 11.1% increase in the number of tornadoes. 270

<sup>271</sup> Changes to the expected number of tornadoes given changes in the environmental variables <sup>272</sup> have a large impact on the probability distribution of counts conditional on the cluster area. The <sup>273</sup> negative binomial distribution for the number of tornadoes *T* with an expected number of tornadoes <sup>274</sup>  $\overline{T}$  (obtained from the regression model) has a probability density

$$\Pr(T=k) = \frac{\Gamma(r+k)}{k!\Gamma(r)} \left(\frac{r}{r+\bar{T}}\right)^r \left(\frac{\bar{T}}{r+\bar{T}}\right)^k \quad \text{for } k = 10, 11, 12, \dots,$$
(2)

where r = 1/n and  $\Gamma(z) = \int_0^\infty x^{z-1} e^{-x} dx$  is the gamma function.

For example, on April 12, 2020 the 12 UTC guidance from the SPC convective outlook defined an 276 area with a 10% chance of at least one tornado occurring within 40 km of any location (10% tornado 277 risk). The area of the polygon was approximately 400,000 km<sup>2</sup> (much larger than the average cluster 278 area) centered on Mississippi (Fig. 7). With an area of that size, the model estimates the probability 279 of at least 30 tornadoes for a range of deep-layer shear values and conditional on the amount of 280 CAPE while holding shallow-layer shear at the average value of all clusters (Fig. 8). Given an 281 average amount of shallow-layer shear, a deep-layer shear of  $10 \text{ m s}^{-1}$  and low CAPE (5th percentile 282 value), the model predicts a 17% [9, 26%, UI] chance of at least 30 tornadoes (given a cluster with 283 at least ten tornadoes). In contrast, given a deep-layer shear of 40 m s<sup>-1</sup> and high CAPE (95th 284 percentile value), the model predicts a 65% [(56, 71%), UI] chance of at least 30 tornadoes. There 285 were more than 100 tornadoes on that day. 286

The model quantifies the empirical relationship between CAPE and, independently, shear in terms of a probability distribution on the number of tornadoes. It predicts the expected count given values for the explanatory variables. The negative binomial distribution uses the model's predicted count and the size parameter to generate a distribution of probabilities. For example, the model gives predicted probabilities across a range of CAPE and deep-layer shear values (holding shallow-layer shear at its mean value) that provides a picture of the relationship (Fig. 9). The predicted probabilities of at least 30 tornadoes given an outbreak covering an area of 400,000 km<sup>2</sup>
 increase from low values of both CAPE and shear to high values of both CAPE and shear.

## *c. Sensitivity of the results to cluster definition*

The clustering methodology is taken from Schroder and Elsner (2019) where a sensitivity analysis 296 was performed to examine the reliability of the resulting clusters. They examined various stopping 297 thresholds and determined that a space-time distance of 5000 s best matched the subjectively 298 identified outbreaks of Forbes (2006) with an agreement of 88%. Larger and smaller stopping 299 thresholds resulting in lower agreement percentages. Still this objective method results in clusters 300 with varying levels of tornado density (tornadoes per unit area), which might influence the results 301 of the regression model. For example, Fig. 1 shows that the June 6, 1999 cluster has much lower 302 tornado density compared to the February 5, 2008 cluster. 303

To directly test the sensitivity of our cluster definition, we first correlate the model residuals 304 (observed count minus the predicted rate) for each cluster with the tornado density. The Pearson 305 product-moment correlation coefficient is .02 indicating that tornado density is not a significant 306 factor in the model's ability to predict the conditional tornado counts from the environmental 307 variables. Second, we refit the model using all clusters except the five clusters having the lowest 308 tornado density. The mean absolute error is only marginally improved from 8.6 to 8.4 tornadoes, 309 providing further evidence that model results are not particularly sensitive to the inclusion of 310 clusters with low tornado density. 311

## *d. A model for the number of casualties*

Next we fit a negative binomial regression to the cluster-level casualty counts (direct injuries and deaths) using the same explanatory variables (Table 3) with the exceptions that population (scaled by 100,000 residents) replaces cluster area and *C* (casualty count) replaces *T* (tornado count) as the dependent variable. This is our casualty-count model. We find that CIN is the only variable not significant in the initial model (Table 5). We remove it before fitting a final model.

The in-sample correlation between the observed casualty counts and predicted rates is .43 [(.37, 318 .48), 95% UI (Fig. 10). We test the significance of population and find the model performance 319 decreases by 12% when excluding population from the model. The mean absolute error between 320 the observed counts and expected rates is 39 casualties or 1.3% of the range in observed counts and 321 3.4% of the range in predicted rates. The out-of-sample correlation is .36 and the mean absolute 322 error is 40 casualties. The skill is lower than the skill of the tornado-count model as there is 323 additional uncertainty associated with the number of casualties given a tornado. The lower skill is 324 also a result of the many other factors that can influence casualties including demographic variables 325 and location (see Summary and Conclusions). 326

As expected from the tornado-count model, the number of casualties resulting from a cluster of 327 tornadoes increases with CAPE and with the two bulk shear variables (Table 5) which is consistent 328 with Anderson-Frey and Brooks (2019). Holding all other variables constant, an increase in CAPE 329 of 1000 J kg<sup>-1</sup> results in a 28% increase in the expected number of casualties. An increase in 330 deep-layer bulk shear of 10 m s<sup>-1</sup> results in a 98% increase in the expected number of casualties 331 per cluster and a similar increase in shallow-layer bulk shear results in a 76% increase in the 332 expected number of casualties per cluster, conditional on at least ten tornadoes. Additionally, the 333 model indicates that casualties decrease at a rate of 3.6% per year. This is very likely the result of 334 improvements made by the National Weather Service in warning coordination and dissemination 335 leading to better awareness especially for these large outbreak events. 336

Also, as expected, the number of people in harm's way is a significant explanatory variable for the cluster-level casualty count. The relationship between population and number of casualties is quantified at the tornado-level in Elsner et al. (2018) and Fricker et al. (2017) so we expect the relationship to hold at the cluster level. Here, we are able to compare the influence of shear and CAPE on the probability of casualties as modulated by population (Fig. 11). Model results are shown for three levels of population. The probability of a large number of casualties increases with increasing shear and increasing CAPE, while keeping the other variables at their mean values and year at 2018.

<sup>345</sup> Importantly, we also find that where the cluster occurs has a significant influence on the number <sup>346</sup> of casualties consistent with other studies (Ashley and Strader 2016; Fricker and Elsner 2019). <sup>347</sup> For every one degree north latitude, the casualty rate decreases by 5.5% and for every one degree <sup>348</sup> east longitude the casualty rate increases by 2.9%. Thus, cluster-level casualties are highest over <sup>349</sup> the Southeast. This effect is independent of the number of tornadoes since location was not a <sup>350</sup> significant factor in the tornado-count model. The result is also independent of the number of <sup>361</sup> people in harm's way since population is included as an exploratory variable in the model.

To visualize the difference of the combined effects of latitude and longitude on the difference in 352 the probability of many casualties, we plot modeled casualty probabilities (at least 25) as a function 353 of CAPE and deep-layer shear for two *hypothetical* outbreaks that are the same in every way except 354 one outbreak is centered on Sioux City, Iowa ( $42.5^{\circ}$  N,  $96.4^{\circ}$  W), and the other is centered on 355 Birmingham, Alabama (33.5° N, 86.8° W) (Fig. 12). The modeled probabilities are lowest (around 356 5%) for low CAPE and shear values and highest (above 30%) for high CAPE and shear values. 357 The difference in modeled probabilities across these two locations peaks at about +12 percentage 358 points for high CAPE and high shear regimes when the outbreak is centered on Birmingham. 359

## **4.** Summary and conclusions

Estimating characteristics of severe weather outbreaks (e.g., tornado and casualty counts) is 361 challenging but important. Forecasters use a combination of numerical weather prediction and 362 empirical guidance to outline areas of severe convective weather. Here, we demonstrate a statistical 363 regression model that can take advantage of the large sample of independent tornado 'outbreaks' as 364 a way to statistically explain the number of tornadoes and the number of casualties in a cluster of at 365 least ten tornadoes. We fit negative binomial regressions to tornado characteristics aggregated to 366 the level of tornado clusters where a cluster is a space-time group of at least ten tornadoes occurring 367 between 12 UTC and 12 UTC over the period 1994–2018. The number of tornadoes in each cluster 368 is the response variable in the tornado-count model, and the number of casualties (deaths plus 369 injuries) is the response variable in the casualty-count model. Environmental explanatory variables 370 for the models are extracted from reanalysis data representing conditions before the occurrence of 371 the first tornado in the cluster consistent with Schroder and Elsner (2019). Additional explanatory 372 variables include cluster area, population, location, and year. 373

The estimated tornado rates, conditional on there being at least ten tornadoes, explain 59% of the 374 observed tornado counts in-sample, and the estimated casualty rates explain 43% of the observed 375 casualty counts in-sample. Because of the large sample size, the out-of-sample skill is lower but 376 still useful. The models show that a 1000 J kg<sup>-1</sup> increase in CAPE results in a 4.7% increase in 377 the expected number of tornadoes conditional on at least ten tornadoes and a 28% increase in the 378 expected number of casualties, holding the other variables constant. The models further show that 379 a 10 m s<sup>-1</sup> increase in deep-layer bulk shear results in a 13% increase in the expected number of 380 tornadoes and a 98% increase in the expected number of casualties, holding the other variables 381 constant. This research is consistent with Anderson-Frey and Brooks (2019) who found the number 382

<sup>383</sup> of tornadoes and casualties to increase with both CAPE and shear. This study quantifies these <sup>384</sup> increases. The casualty-count model also shows a significant decline in the number of casualties <sup>385</sup> at a rate of 3.6% per year. Casualty rates depend on where the outbreak occurs with more deaths <sup>386</sup> and injuries, on average, over the Southeast, controlling for the other variables; a result that is <sup>387</sup> consistent with the recent work of Fricker and Elsner (2019) and Biddle et al. (2020).

Some of the unexplained variability in cluster-level tornado counts (and casualty counts) arises 388 from the uncertainty associated with the preferred storm mode and the evolution of meso-scale 389 convective systems, neither of which are captured by a single maximum value in the variable space 390 of CAPE and shear. The counts are also limited by the quality of the NARR data. The NARR 391 tends to unrealistically favor tornado environments during specific convective setups (Gensini and 392 Ashley 2011; Gensini et al. 2014; Allen et al. 2015a). Additionally, we use only the maximum 393 values (minimum for CIN) of the environmental variables which may limit the representation of 394 the cluster environment. Also, outbreaks associated with tropical cyclones likely add a bit of noise 395 to both models since the number of tornadoes is sensitive to the extent and location of convective 396 bursts within overall evolution of the land-falling storm. 397

The casualty-count model would be improved by including a skillful estimate of the number 398 of tornadoes. Indeed in a perfect-prognostic setting, where we know the number of tornadoes in 399 the outbreak, the out-of-sample correlation between the observed number of casualties and the 400 modeled estimated rate of casualties increases to .79. Further, although our approach to extracting 401 signal from noise in the tornado dataset is sound, exclusive focus on clusters with at least ten 402 tornadoes is a type of selection bias meaning that the sample of data used to fit the model does not 403 represent the population of all outbreaks, which limits what we can say in general about the effect 404 of convective environments on the probability distribution of casualty counts. 405

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The tornado-count model can be modified to provide guidance to forecasters given a convective 406 outlook that highlights an area of elevated threat for tornadoes and a prediction of CAPE and 407 shear across the threat area. The model needs to be calibrated using threat polygons (not cluster 408 areas as was done here) and include *predicted* environmental values, but the same model equation 409 can be used to provide a forecast probability distribution on the number of tornadoes. Further, a 410 numerical convolution of this probability distribution with a forecast probability distribution for 411 each EF-rating category (Elsner and Schroder 2019) will result in a prediction of the expected 412 number of counts by category as well as the associated uncertainties. 413

The casualty-count model can be employed in a research setting to help better understand the socioeconomic, demographic, and communication factors that make some communities particularly vulnerable to deaths and injuries (Dixon and Moore 2012; Senkbeil et al. 2013; Klockow et al. 2014; Fricker and Elsner 2019). Work along this line has been done at the individual tornado level by identifying unusually devastating events (Fricker and Elsner 2019), but scaling this type of analysis to the cluster-level to identify unusually devastating outbreaks might provide additional insights.

Finally, it is possible that the models could be improved by including nonlinear effects. One type 421 of non-linearity is interaction where the effect of CAPE on casualties is modulated by shear, for 422 example. However, interaction effects usually must be specified without reference to the data, so 423 additional research on this is needed. The model could also be improved by including variables 424 that represent other environmental factors, convective modes, and efficiency of tornado production. 425 The models also might be improved by adjusting the threshold definition of a cluster. Increasing 426 the threshold on the tornado-count model from 10 to 14 decreases the sample size to 505 clusters 427 and reduces the effect sizes on CAPE and shear by around 25%. Decreasing the threshold from 10 428 to 6 increases the sample size and, thus, reduces the standard error assuming the effect size stays 429

the same. A casualty-count model might also be improved by relaxing the assumption that the
numbers of people injured or killed are independent. Casualty counts are typically not independent
at the household level where multiple people live under the same roof. In this case a better model
might include a zero-inflated count process.

Acknowledgments. The negative binomial regression models in this paper were implemented with
the glm.nb function from the MASS R package (Venables and Ripley 2002). Graphics were made
with the ggplot2 framework (Wickham 2017). The code and data to fit all the models is available
on GitHub (https://github.com/jelsner/cape-shear).

# 438 **References**

439	Allen, J. T., M. K. Tippett, and A. H. Sobel, 2015a: An empirical model relating U.S. monthly hail
440	occurrence to large-scale meteorological environment. Journal of Advances in Modeling Earth
441	Systems, 7 (1), 226–243, doi:10.1002/2014MS000397, URL https://agupubs.onlinelibrary.wiley.
442	com/doi/abs/10.1002/2014MS000397.
443	Allen, J. T., M. K. Tippett, and A. H. Sobel, 2015b: Influence of the El Niño/Southern Oscillation
444	on tornado and hail frequency in the United States. Nature Geosciences, 8, 278–283.
445	Anderson-Frey, A. K., and H. Brooks, 2019: Tornado fatalities: An environmental perspective.
446	Weather and Forecasting, 34 (6), 1999–2015, doi:10.1175/waf-d-19-0119.1, URL https://doi.
447	org/10.1175/waf-d-19-0119.1.
448	Anderson-Frey, A. K., Y. P. Richardson, A. R. Dean, R. L. Thompson, and B. T. Smith, 2018:
449	Near-storm environments of outbreak and isolated tornadoes. Weather and Forecasting, 33 (5),
450	1397-1412, doi:10.1175/WAF-D-18-0057.1, URL https://doi.org/10.1175/WAF-D-18-0057.1,
451	https://doi.org/10.1175/WAF-D-18-0057.1.
452	Ashley, W. S., and S. M. Strader, 2016: Recipe for disaster: How the dynamic ingredients of risk and
453	exposure are changing the tornado disaster landscape. Bulletin of the American Meteorological
454	Society, <b>97</b> , 767–786.
455	Biddle, M. D., R. P. Brown, C. A. Doswell, and D. R. Legates, 2020: Regional differences in the
456	human toll from tornadoes: A new look at an old idea. Weather, Climate, and Society, 12 (4),
457	815-825, doi:10.1175/wcas-d-19-0051.1, URL https://doi.org/10.1175/wcas-d-19-0051.1.
458	Brooks, H. E., G. W. Carbin, and P. T. Marsh, 2014: Increased variability of tornado occur-

rence in the United States. *Science*, **346** (**6207**), 349–352, doi:10.1126/science.1257460, URL

22

https://science.sciencemag.org/content/346/6207/349, https://science.sciencemag.org/content/ 460 346/6207/349.full.pdf. 461

462	Brooks, H. E., C. A. Doswell , and J. Cooper, 1994: On the environments of tornadic and
463	nontornadic mesocyclones. Weather and Forecasting, 9, 606–618, doi:10.1175/1520-0434.
464	Cohen, A. E., J. B. Cohen, R. L. Thompson, and B. T. Smith, 2018: Simulating tornado probability
465	and tornado wind speed based on statistical models. Weather and Forecasting, 33 (4), 1099–1108,
466	doi:10.1175/waf-d-17-0170.1, URL https://doi.org/10.1175/waf-d-17-0170.1.
467	Concannon, P., H. E. Brooks, and C. A. Doswell , 2000: Climatological risk of strong and violent
468	tornadoes in the United States. Second Conference on Environmental Applications.
469	Dean, A. R., 2010: P2.19 An analysis of clustered tornado events. 25th Conference on Severe Local
470	Storms.
471	Dixon, P. G., A. E. Mercer, K. Grala, and W. H. Cooke, 2014: Objective identification of tornado
472	seasons and ideal spatial smoothing radii. <i>Earth Interactions</i> , <b>18</b> , 1–15.
473	Dixon, R. W., and T. W. Moore, 2012: Tornado vulnerability in Texas. Weather, Climate, and

*Society*, **4**, 59–68.

474

- Doswell, C. A., R. Edwards, R. L. Thompson, J. A. Hart, and K. C. Crosbie, 2006: A simple and 475 flexible method for ranking severe weather events. Weather and Forecasting, 21 (6), 939–951, 476 doi:10.1175/waf959.1, URL https://doi.org/10.1175/waf959.1. 477
- Elsner, J. B., S. C. Elsner, and T. H. Jagger, 2015: The increasing efficiency of tornado days in the 478 United States. Climate Dynamics, 45 (3-4), 651–659. 479

- Elsner, J. B., T. Fricker, and W. D. Berry, 2018: A model for U.S. tornado casualties involving
   interaction between damage path estimates of population density and energy dissipation. *Journal of Applied Meteorology and Climatology*, **57**, 2035–2046.
- Elsner, J. B., and C. P. Schmertmann, 1994: Assessing forecast skill through cross validation. *Weather and Forecasting*, **9** (**4**), 619–624.
- Elsner, J. B., and Z. Schroder, 2019: Tornado damage ratings estimated with cumulative logistic regression. *Journal of Applied Meteorology and Climatology*, **58** (**12**), 2733–2741, doi:10.1175/

<sub>487</sub> jamc-d-19-0178.1, URL https://doi.org/10.1175/jamc-d-19-0178.1.

- Forbes, G. S., 2006: Meteorological aspects of high-impact tornado outbreaks. Preprints, 22nd
   *Conf. on Severe Local Storms*, Hyannis, MA, Amer. Meteor. Soc., 1–12.
- 490 Fricker, T., and J. B. Elsner, 2019: Unusually devastating tornadoes in the United States:
- <sup>491</sup> 1995–2016. Annals of the American Association of Geographers, **110** (**3**), 724–738, doi:
- <sup>492</sup> 10.1080/24694452.2019.1638753, URL https://doi.org/10.1080/24694452.2019.1638753.
- Fricker, T., J. B. Elsner, and T. H. Jagger, 2017: Population and energy elasticity of tornado casualties. *Geophysical Research Letters*, **44**, 3941–3949, doi:10.1002/2017GL073093.
- Fujita, T. T., 1981: Tornadoes and downbursts in the context of generalized planetary scales. *J. Atmos. Sci.*, **38**, 1511–1534.
- Galway, J. G., 1977: Some climatological aspects of tornado outbreaks. *Monthly Weather Review*,
   **105**, 477–484.
- Gensini, V., and W. Ashley, 2011: Climatology of potentially severe convective environments from
   the North American Regional Reanalysis. *Electronic Journal of Severe Storms Meteorology*, 6,
   1–40, doi:10.1038/s41612-018-0048-2.

- Gensini, V. A., T. L. Mote, and H. E. Brooks, 2014: Severe-thunderstorm reanalysis environments
   and collocated radiosonde observations. *Journal of Applied Meteorology and Climatology*,
   53 (3), 742–751, doi:10.1175/jamc-d-13-0263.1, URL https://doi.org/10.1175/jamc-d-13-0263.
   1.
- <sup>506</sup> Heiss, W. H., D. L. McGrew, and D. Sirmans, 1990: NEXRAD: next generation weather radar <sup>507</sup> (WSR-88D). *Microwave Journal*, **33** (1), 79.
- <sup>508</sup> Hilbe, J., 2011: *Negative Binomial Regression*. Cambridge University Press.
- Hill, A. J., G. R. Herman, and R. S. Schumacher, 2020: Forecasting severe weather with random
- <sup>510</sup> forests. *Monthly Weather Review*, doi:10.1175/mwr-d-19-0344.1, URL https://doi.org/10.1175/ <sup>511</sup> mwr-d-19-0344.1.
- Hitchens, N. M., and H. E. Brooks, 2014: Evaluation of the Storm Prediction Center's convective
  outlooks from day 3 through day 1. *Weather and Forecasting*, 29 (5), 1134–1142, doi:10.1175/
  waf-d-13-00132.1, URL https://doi.org/10.1175/waf-d-13-00132.1.
- Jackson, J. D., and M. E. Brown, 2009: Sounding-derived low-level thermodynamic characteristics associated with tornadic and non-tornadic supercell environments in the Southeast United States. *National Weather Digest*, **33**, 16–26.
- Klockow, K. E., R. A. Peppler, and R. A. McPherson, 2014: Tornado folk science in Alabama
- and Mississippi in the 27 April 2011 tornado outbreak. *GeoJournal*, **79** (6), 791–804, doi:

<sup>520</sup> 10.1007/s10708-013-9518-6, URL https://doi.org/10.1007/s10708-013-9518-6.

- Mesinger, F., and Coauthors, 2006: North American Regional Reanalysis. Bulletin of the American
- 522 Meteorological Society, 87 (3), 343–360, doi:10.1175/BAMS-87-3-343, URL https://doi.org/
- <sup>523</sup> 10.1175/BAMS-87-3-343, https://doi.org/10.1175/BAMS-87-3-343.

<sup>524</sup> Moore, T., 2017: On the temporal and spatial characteristics of tornado days in the United States. *Atmospheric Research*, **184**, doi:10.1016/j.atmosres.2016.10.007.

Moore, T. W., R. W. Dixon, and N. J. Sokol, 2016: Tropical cyclone Ivan's tornado cluster
 in the Mid-Atlantic region of the United States on 17–18 September 2004. *Physical Geogra- phy*, **37** (3-4), 210–227, doi:10.1080/02723646.2016.1189299, URL https://doi.org/10.1080/
 02723646.2016.1189299.

- Rasmussen, E. N., and D. O. Blanchard, 1998: A baseline climatology of sounding derived supercell and tornado forecast parameters. *Weather and Forecasting*, 13 (4), 1148–
   1164, doi:10.1175/1520-0434(1998)013<1148:ABCOSD>2.0.CO;2, URL https://doi.org/10.
   1175/1520-0434(1998)013<1148:ABCOSD>2.0.CO;2.
- Schroder, Z., and J. B. Elsner, 2019: Quantifying relationships between environmental factors
   and power dissipation on the most prolific days in the largest tornado "outbreaks". *International Journal of Climatology*, doi:10.1002/joc.6388, URL https://doi.org/10.1002/joc.6388.
- Senkbeil, J. C., D. A. Scott, P. Guinazu-Walker, and M. S. Rockman, 2013: Ethnic and racial
   differences in tornado hazard perception, preparedness, and shelter lead time in Tuscaloosa. *The Professional Geographer*, 66 (4), 610–620, doi:10.1080/00330124.2013.826562, URL https:
   //doi.org/10.1080/00330124.2013.826562.
- Shafer, C. M., and C. A. Doswell, 2011: Using kernel density estimation to identify, rank, and
   classify severe weather outbreak events. *Electronic Journal of Severe Storms Meteorology*, 6,
   1–28.
- Steiner, E., 2019: Spatial history project. Center for Spatial and Textual Analysis, Stanford Uni versity, URL http://web.stanford.edu/group/spatialhistory/cgi-bin/site/index.php.

546	Thompson, R. L., R. Edwards, J. A. Hart, K. L. Elmore, and P. Markowski, 2003: Close proxim-
547	ity soundings within supercell environments obtained from the Rapid Update Cycle. Weather
548	and Forecasting, <b>18</b> ( <b>6</b> ), 1243–1261, doi:10.1175/1520-0434(2003)018<1243:cpswse>2.0.co;2,
549	URL https://doi.org/10.1175/1520-0434(2003)018<1243:cpswse>2.0.co;2.
550	Thompson, R. L., and Coauthors, 2017: Tornado damage rating probabilities derived from WSR-
551	88D data. Weather and Forecasting, 32 (4), 1509–1528, doi:10.1175/waf-d-17-0004.1, URL
552	https://doi.org/10.1175/waf-d-17-0004.1.
553	Tippett, M. K., C. Lepore, and J. E. Cohen, 2016: More tornadoes in the most extreme U.S.
554	tornado outbreaks. Science, 354 (6318), 1419–1423, doi:10.1126/science.aah7393, URL https:
555	//doi.org/10.1126/science.aah7393.
556	Tippett, M. K., A. H. Sobel, and S. J. Camargo, 2012: Association of U.S. tornado occurrence
557	with monthly environmental parameters. Geophysical Research Letters, 39, L02 801.
558	Venables, W. N., and B. D. Ripley, 2002: Modern Applied Statistics with S. 4th ed., Springer, New
559	York, URL http://www.stats.ox.ac.uk/pub/MASS4, iSBN 0-387-95457-0.
560	Wickham, H., 2017: tidyverse: Easily Install and Load 'Tidyverse' Packages. URL https://CRAN.
561	R-project.org/package=tidyverse, r package version 1.1.1.

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TABLE 1. Cluster statistics by time of day. Each cluster is categorized by the closest three-hour time (defined by the NARR data) prior to the first tornado.

Time of Day (UTC)	Number of Clusters	Number of Tornadoes	Average Tornadoes Per Cluster	Average Duration (hours)
00	33	523	15.8	6.1
03	5	67	13.4	6.4
06	2	23	11.5	3.2
12	145	3598	12.1	14.0
15	124	3222	26.0	11.5
18	249	5220	21.0	8.4
21	210	4416	21.0	7.0

TABLE 2. Correlation matrix of environmental variables considered in this study. Dew-point temperature (DEW), specific humidity (SH), and storm relative helicity (HLCY). Only CAPE, CIN, DLBS, and SLBS are used as explanatory variables in the models.

	CAPE	CIN	DLBS	SLBS	HLCY	DEW	SH
CAPE	1.00						
CIN	-0.07	1.00					
DLBS	-0.03	-0.29	1.00				
SLBS	-0.37	-0.24	0.49	1.00			
HLCY	-0.22	-0.30	0.58	0.76	1.00		
DEW	0.56	0.00	-0.08	0.02	-0.08	1.00	
SH	0.64	0.00	-0.12	-0.08	-0.13	0.98	1.00

TABLE 3. Variables considered in the regression models. Values include the range and average across the 768
 tornado clusters.

Variable	Abbreviation	Range	Average				
Explanatory Variables							
Convective Available Potential Energy [J kg <sup>-1</sup> ]	CAPE	[0, 6530]	2225				
Convective Inhibition [J kg <sup>-1</sup> ]	CIN	[-668, 0]	-114				
Deep-Layer Bulk Shear [m s <sup>-1</sup> ]	DLBS	[5.6, 48]	27.5				
Shallow-Layer Bulk Shear [m s <sup>-1</sup> ]	SLBS	[1.1, 33.8]	15.0				
Latitude [° N]	$\phi$	[27.12, 48.97]	37.20				
Longitude [° E]	λ	[-109.9 -72.88]	-92.16				
Cluster Area [km <sup>2</sup> ]	A	[361, 1,064,337]	167,990				
Population [No. of People]	Р	[0, 38,226,946]	3,387,259				
Year	Y	[1994, 2018]	2006				
Response Variables							
Number of Tornadoes	Т	[10, 173]	22.2				
Number of Casualties (injuries plus deaths)	С	[0, 3,069]	29.9				

TABLE 4. Coefficients in the tornado-count models. The size parameter (*n*) is  $6.27 \pm .393$  (standard error) for the initial model  $6.25 \pm .392$  (standard error) for the final model.

Coefficient	Estimate	S.E.	z value	Pr(> z )		
	Initial Model					
$eta_0$	4.5489	4.7662	0.9540	0.3399		
$eta_A$	0.0146	0.0011	12.80	< 0.0001		
$eta_{\phi}$	-0.0051	0.0043	-1.17	0.2427		
$\beta_{\lambda}$	-0.0028	0.0031	-0.917	0.3594		
$\beta_Y$	-0.0012	0.0024	-0.515	0.6068		
$\beta_{CAPE}$	0.0452	0.0153	2.96	0.0031		
$\beta_{CIN}$	-0.0110	0.0189	-0.581	0.5612		
$\beta_{DLBS}$	0.1256	0.0292	4.30	< 0.0001		
$\beta_{SLBS}$	0.1059	0.0355	2.98	0.0029		
		Final	Model			
$eta_0$	2.1779	0.0817	26.65	< 0.0001		
$eta_A$	0.0149	0.0011	13.85	< 0.0001		
$\beta_{CAPE}$	0.0459	0.0146	3.13	0.0017		
$\beta_{DLBS}$	0.1254	0.0288	4.35	< 0.0001		
$\beta_{SLBS}$	0.1054	0.0314	3.35	0.0008		

Coefficient	Estimate	S.E.	z value	Pr(> z )		
	Initial Model					
$eta_0$	76.6908	20.7430	3.70	0.0002		
$\beta_P$	0.0122	0.0019	6.51	< 0.0001		
$eta_{\phi}$	-0.0561	0.0187	-3.00	0.0027		
$\beta_{\lambda}$	0.0284	0.0136	2.09	0.0363		
$\beta_Y$	-0.0364	0.0103	-3.52	0.0004		
$\beta_{CAPE}$	0.2436	0.0643	3.79	0.0002		
$\beta_{CIN}$	0.0052	0.0802	0.07	0.9479		
$\beta_{DLBS}$	0.6853	0.1262	5.43	< 0.0001		
$\beta_{SLBS}$	0.5650	0.1534	3.68	0.0002		
		Final	Model			
$eta_0$	76.7677	20.6902	3.71	0.0002		
$\beta_P$	0.0122	0.0018	6.67	0.0000		
$eta_{\phi}$	-0.0563	0.0186	-3.02	0.0025		
$\beta_{\lambda}$	0.0287	0.0130	2.20	0.0277		
$\beta_Y$	-0.0364	0.0103	-3.53	0.0004		
$\beta_{CAPE}$	0.2440	0.0643	3.79	0.0001		
$\beta_{DLBS}$	0.6833	0.1253	5.45	0.0000		
$\beta_{SLBS}$	0.5631	0.1504	3.74	0.0002		

TABLE 5. Coefficients in the casualty-county models. The size parameter (*n*) is  $.261 \pm .014$  (standard error) for the initial and final models.

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613 614 615 616	Fig. 8.	Estimated probability of at least 30 tornadoes given an outbreak of at least ten tornadoes and the regression model. The predicted count from the model is a parameter in a negative binomial distribution with cluster area set at 400,000 km <sup>2</sup> and shallow-level bulk shear is set to its mean value.	43
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620 621 622 623	Fig. 10.	Observed cluster-level casualty counts versus predicted rates from a negative binomial re- gression. Clusters without casualties are plotted at the far left. The thin black line is the line of best fit. The thick line is the slope of the model indicating the relationship between the observed and predicted tornado counts and the associated standard error.	45
624 625 626	Fig. 11.	Probability of at least 50 tornado casualties as a function of deep-layer bulk shear (left panel) and CAPE (right panel) and modulated by the number of people in harm's way. The other variables are set at their mean values and year is set at 2018.	46

627	Fig. 12.	Probability of at least 25 tornado casualties as a function of deep-layer bulk shear and CAPE	
628		and modulated by location for two hypothetical outbreaks, one centered over Sioux City,	
629		Iowa, and the other centered over Birmingham, Alabama. The shallow-layer bulk shear is	
630		set to its mean value, year is set to 2018, and population is set to 4M.	47

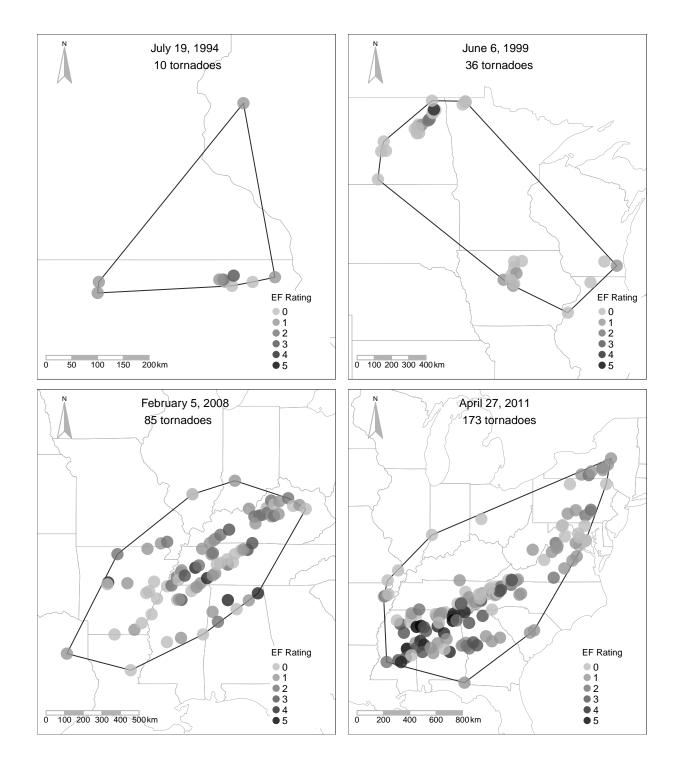


FIG. 1. Example tornado clusters. Each point is the tornadogenesis location shaded by EF rating. The black line is the spatial extent of the tornadoes occurring on that convective day and is defined by the minimum convex hull encompassing the set of locations.

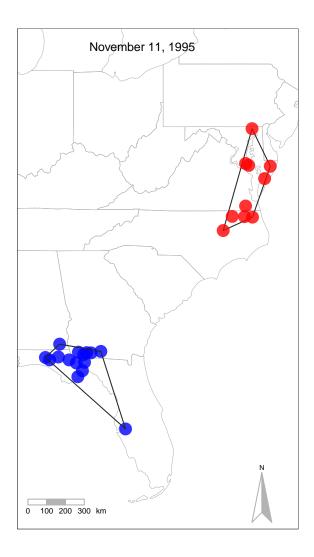


FIG. 2. Example of multiple clusters on a single convective day. Each point is a tornado genesis location. The black line is the spatial extent of the tornadoes for each cluster and is defined by the minimum convex hull encompassing the set of locations

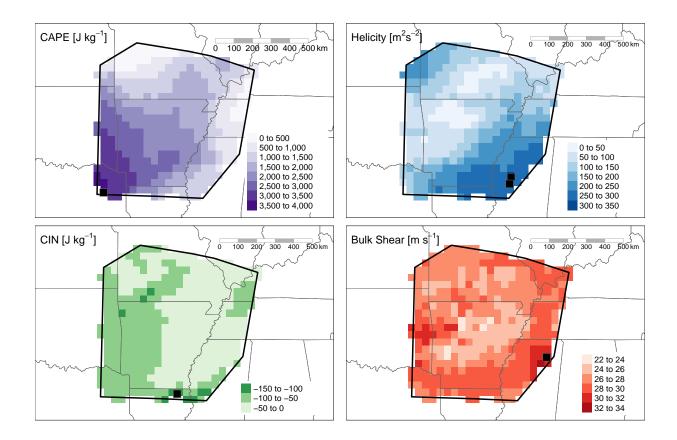


FIG. 3. Example of the environmental factors for the May 6, 2003. The black line is the spatial extent of the cluster on that convective day. Shading represents the intensity of the environment. CAPE is purple, helicity is blue, CIN is green, and deep layer shear is red. The black square is the location of the maximum value for the environmental factor.

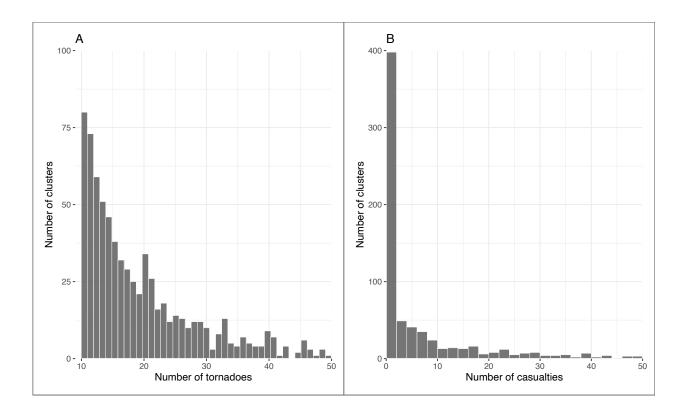


FIG. 4. Histograms of the number of clusters by number of tornadoes (A) and number of clusters by number of casualties (B). The histograms are right-truncated at 50 to show detail on the left side of the distributions. Only clusters with at least ten tornadoes are considered in this study.

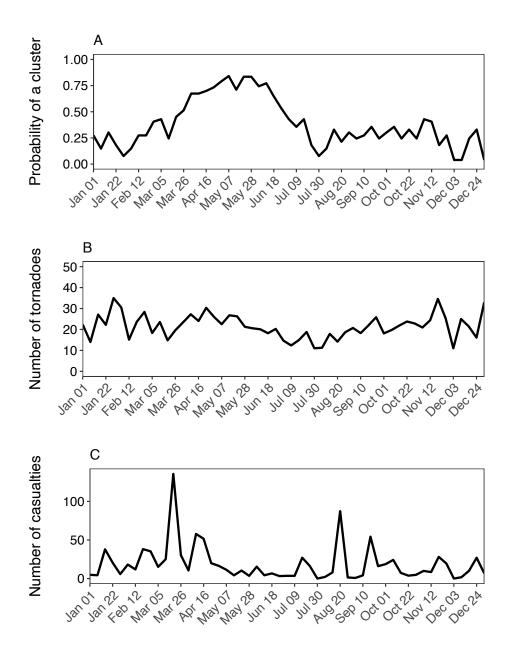


FIG. 5. Probability of a cluster (A), average number of tornadoes per cluster (B), and average number of casualties per 1 000 000 people per cluster (C) by week of the year.

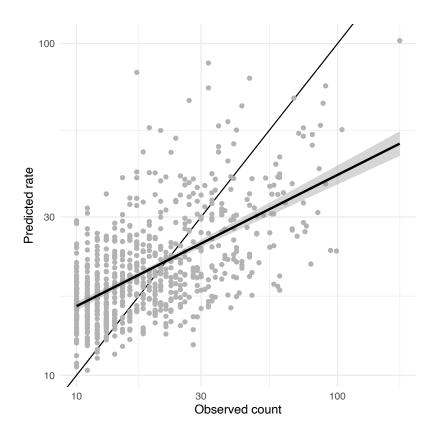


FIG. 6. Observed cluster-level tornado counts versus predicted rates from a negative binomial regression. The thin black line is the line of best fit. The thick line is the slope of the model indicating the relationship between the observed and predicted tornado counts and the associated standard error.

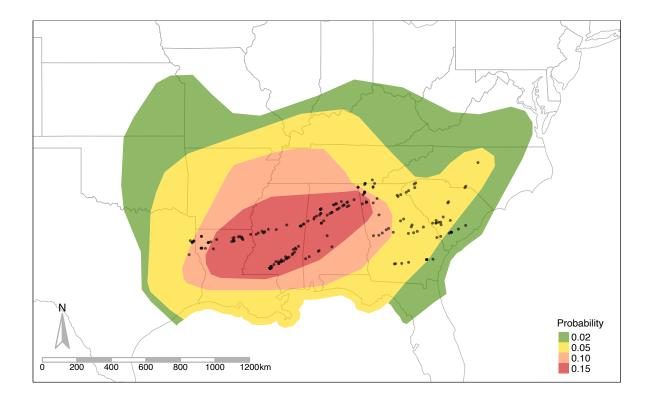


FIG. 7. Convective outlook issued by the Storm Prediction Center at 12 UTC on April 12, 2020 and the locations of tornado reports over the 24-hr period starting at that time. The outlook category numbers indicate the chance of observing severe weather within 40 km of any location. The convective outlook shapefiles are from www.spc.noaa.gov/cgi-bin-spc/getacrange.pl?date0=20200412&date1=20200412 and the tornado reports are from www.spc.noaa.gov/climo/reports/200412\_rpts.html.

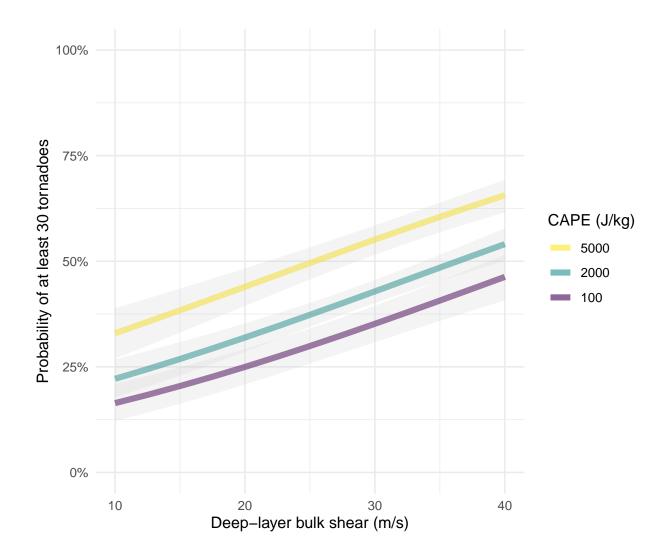


FIG. 8. Estimated probability of at least 30 tornadoes given an outbreak of at least ten tornadoes and the regression model. The predicted count from the model is a parameter in a negative binomial distribution with cluster area set at 400,000 km<sup>2</sup> and shallow-level bulk shear is set to its mean value.

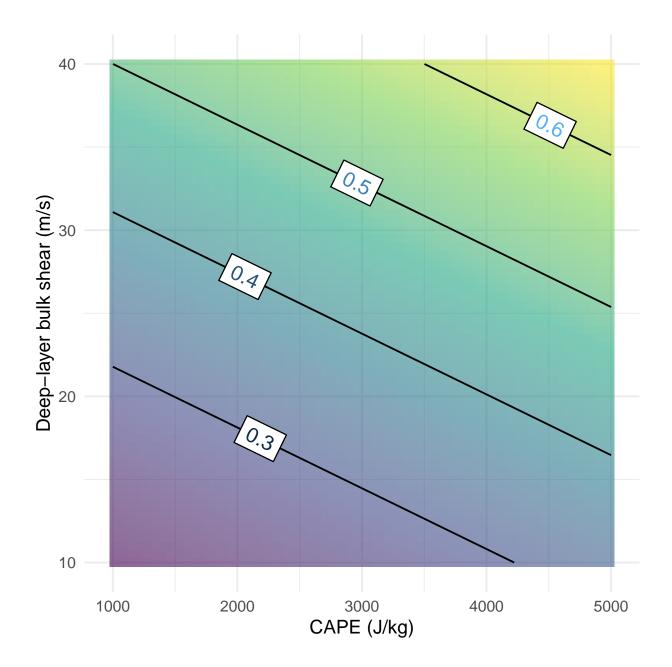


FIG. 9. Estimated probability of at least 30 tornadoes given an outbreak of at least ten tornadoes and the regression model across a range of CAPE and deep-layer bulk shear values holding the shallow-layer bulk shear at a mean value.

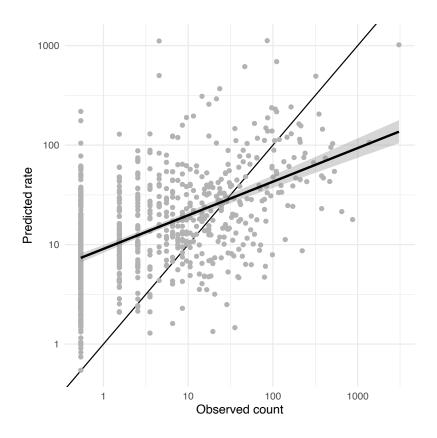


FIG. 10. Observed cluster-level casualty counts versus predicted rates from a negative binomial regression. Clusters without casualties are plotted at the far left. The thin black line is the line of best fit. The thick line is the slope of the model indicating the relationship between the observed and predicted tornado counts and the associated standard error.

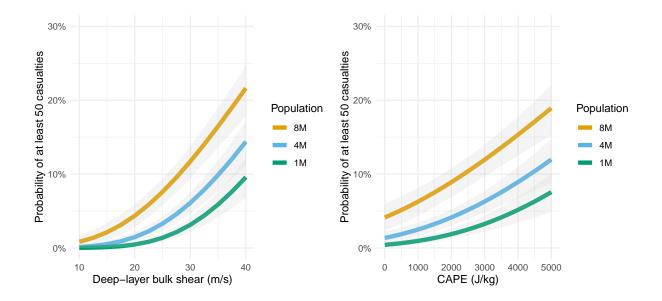


FIG. 11. Probability of at least 50 tornado casualties as a function of deep-layer bulk shear (left panel) and CAPE (right panel) and modulated by the number of people in harm's way. The other variables are set at their mean values and year is set at 2018.

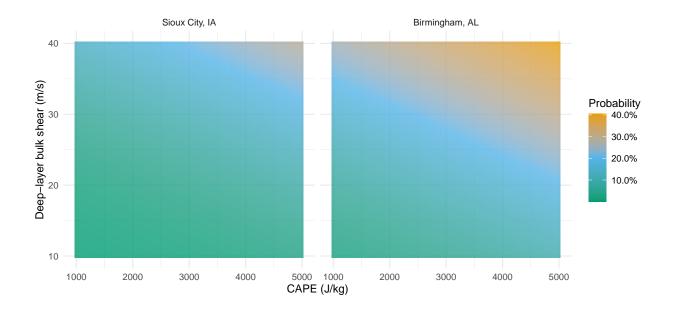


FIG. 12. Probability of at least 25 tornado casualties as a function of deep-layer bulk shear and CAPE and modulated by location for two *hypothetical* outbreaks, one centered over Sioux City, Iowa, and the other centered over Birmingham, Alabama. The shallow-layer bulk shear is set to its mean value, year is set to 2018, and population is set to 4M.