Lateral variability in strain along a mass-transport complex (MTC) toewall: a case study from the Makassar Strait, offshore Indonesia

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Abstract: Contractional features characterise the toe domain of mass-transport complexes (MTCs). Their frontal geometry is typically classified as frontally-confined or frontally-emergent. However, it remains unclear how frontal emplacement style and contractional strain within an MTC can vary along strike. We use bathymetry and 3D seismic reflection data to investigate lateral variability of frontal emplacement and strain within the toe domain of a recent MTC (the Haya Slide) in the Makassar Strait. The slide originated from an anticline flank collapse, and the toe domain is characterised by a radial fold-and-thrust belt that reflects southwestwards emplacement. The degree of frontal confinement and overall toe domain structural style changes laterally, from frontally-confined in the S (associated with a deeply-incised, *c.* 200 mbsf, planar basal shear surface), to frontally-emergent in the W (associated with a shallowly incised, *c.* 120 mbsf, NE-dipping, *c.* 3°, basal shear surface). Quantitative strain analysis shows *c.* 8-14% shortening in the toe domain and that strain within the fold-and-thrust belt varies laterally. This study shows that even minor horizontal translation of MTCs (*c.* 1 km) can drive marked lateral variability in frontal geometry and strain within the failed body, which may influence their seal potential in petroleum and carbon storage systems.

Mass-transport complexes (MTCs) are the deposits of creep, slide, slump, and debris flow processes (e.g. Dott 1963; Nardin et al. 1979; Nemec 1991; Moscardelli & Wood 2008; Posamentier & Martinsen 2011). MTC emplacement can cause major geohazards for offshore infrastructures and coastal communities (e.g. Tappin et al. 2001; Vanneste et al. 2013; Takagi et al. 2019) and can be an important component of a functional petroleum system (e.g. Weimer & Shipp 2004). For example, MTCs can provide seals for hydrocarbon accumulations (Algar et al. 2011; Omeru 2014; Cardona et al. 2016) and, less commonly, may act as reservoirs (Sawyer et al. 2007; Shanmugam 2012; Arfai et al. 2016). In particular, their seal potential depends on a combination of the lithology, external geometry and internal structural heterogeneity of the emplaced mass, which are all influenced by emplacement processes (e.g. Alves et al. 2014). Thus, it is important to understand their transport processes to assess their seal potential in a petroleum system.

The nature of the failed mass in the vicinity of the toewall defines two frontal geometrical types (Frey-Martínez et al. 2006): (i) frontally-confined types represent toewall trapping of a failed mass from further downdip translation, and (ii) frontally-emergent types reflect a failed mass that extends above and beyond the toewall with further translation onto the adjacent seabed. In some cases, both styles can develop within a single mass-transport event (Moernaut & De Batist 2011; Armandita et al. 2015; Clare et al. 2018). The seismic expression of both frontal termination types are well-known (Trincardi & Argnani 1990; Huvenne et al. 2002; Lastras et al. 2004; Joanne et al. 2013), but the processes occurring in the toe domain remain poorly constrained (e.g. evolution of the basal shear surface prior to termination at the toewall). Outcrop studies have provided detailed insights on processes in the toe domain, but a full 3D analysis is hindered by limited exposure extent (Martinsen & Bakken 1990; Van Der Merwe et al. 2011; Sobiesiak et al. 2016; Cardona et al. 2020). Furthermore, very few studies have attempted to balance extensional and contractional strains across the entire body of an MTC (e.g. Bull & Cartwright 2019; Steventon et al. 2019). Likewise, the way in which strain varies along-strike within an MTC remains poorly understood.

Here, we use high-resolution multibeam bathymetry and high-quality 3D seismic reflection data to study the Haya Slide (hereafter the 'slide'), offshore western Sulawesi (Indonesia). This dataset demonstrates how frontal toewall style can change laterally during emplacement of a single mass-transport event. The bathymetry data capture the seabed expression of both the headwall and toe domains of this slide, while the 3D seismic reflection data only image the toe domain, which is the focus of this study (Fig. 1). The seismic image quality and use of seismic attributes enable us to characterise intra-MTC strain in great detail. Our specific aims are to: (i) evaluate kinematic indicators and reconstruct transport processes of the slide, (ii) assess lateral variability of the slide's frontal geometry and infer its controlling factors, (iii) quantitatively examine along-strike changes of intra-MTC strain, and (iv) discuss how lateral variations in strain may impact the seal potential of MTCs.

GEOLOGICAL SETTING

The Makassar Strait is situated within an earthquake-prone region, where four major plates interact (the Eurasia, Indo-Australia, Philippine Sea, and Pacific plates; Fig. 1a) (Daly et al. 1991). A strong southwards-flowing current, the Indonesia Throughflow (ITF), presently carries water masses through the strait at a relatively high velocity (i.e. 1 m/s; Mayer & Damm 2012), from the Pacific Ocean to the Indian Ocean. Brackenridge et al. (2020) suggest that the ITF preconditions the slopes bounding the straits to fail, whereas earthquakes in this seismic-prone region may act as a trigger mechanism. More specifically, the ITF transports a high suspended sediment load southward from the Mahakam Delta, causing relatively rapid deposition and steepening of the continental slope along the western margin of the strait, which results in (i) slope oversteepening, and (ii) high pore-fluid pressures (Brackenridge et al. 2020). Such preconditioning factors for slope failure are consistent with the unusually large number of near-seabed MTCs (Pleistocene to Recent), which range in size from 5 to >600 km^3 (Brackenridge et al. 2020).

The Makassar Strait is located between the islands of Sulawesi and Kalimantan (Borneo) and is divided into the North and South Makassar basins (Fig. 1b). The water depth along the strait is 200-2000 m

(Guntoro 1999), with (i) a relatively broad, accretionary shelfal area along the western margin (including the Mahakam Delta; e.g. Allen & Chambers 1998), and (ii) a narrower and steeper shelf along the eastern margin, which is more tectonically active and bounded by three fold-thrust belts, namely the Northern (NSP), Central (CSP) and Southern (SSP) structural provinces (Puspita et al. 2005). These two marginal areas are the sources of the MTCs transported into the basins (Fig. 1c). The two basins are connected by the deep (*c.* 2000 m) and narrow (*c.* 45 km-wide) Labani Channel, and are cut by major structural features, such as the Palu-Koro and Paternoster transform fault zones (Cloke et al. 1999) (Fig. 1b). We here focus on the Haya Slide (Fig. 1d), which is a shallowly buried MTC with only a thin (<8 m) veneer of modern sediment and a clear present-day seabed expression. The slide is located at the southern end of the Labani Channel, close to the southern margin of the SSP (Fig. 1b).

DATA SET AND METHODOLOGY

The study is based primarily on bathymetry, 3D seismic reflection and well data (Fig. 1b and d). TGS provided the multibeam echosounder bathymetry data (TGS_Pat survey), which covers an area of *c.* 20,000 km^2. Lateral resolution of these data is 25 x 25 m and geomorphic features are enhanced by a shaded relief map with 0^o azimuth and 45^o angle. Core descriptions of near-seabed sediments (*c.* 3-7 mbsf) are also available (i.e. TGS009 and TGS194, see Fig. 1b). Although none of these cores directly sample the Haya Slide, they enable the likely lithology of the slide to be inferred.

The post-stack time-migrated (PSTM) 3D seismic reflection and exploration well data (see Fig. 1b) are provided by the Information and Data Centre, Ministry of Energy and Mineral Resources (PUSDATIN ESDM), Indonesia. The seismic reflection data cover an area of 1598 km^2**,** with a bin spacing of 25 m x 12.5 m (inline x crossline) and a dominant frequency of 50 Hz at the base of the Haya Slide (*c.* 200 mbsf). We estimate that the spatial resolution of the seismic data, given an average velocity of the sedimentary package of interest derived from the wells (1495 m/s), is *c.* 7 m. The average velocity of the near-seabed sediments is relatively low, likely due to the high water content. Similar values are obtained for near-seabed, deep-water sediments penetrated in the South Makassar MTC area, which is located *c.* 135 km to the SW of our study area (Armandita et al. 2015). The 3D seismic data are zero-

phase with SEG normal polarity with an increase in acoustic impedance expressed as a positive amplitude. The two wells (XR-1 and XS-1) do not penetrate the Haya Slide, and there are no drill cuttings data available, even within the general stratigraphic interval containing the slide. However, the correlation of the basal shear surface to the wells ('detachment level' in Fig. 1d) enables the velocity of the sedimentary package containing the slide to be inferred. Using these data allows the conversion of measured vertical distances in time (ms TWT) to depth (m).

The bathymetry data allow delineation of the external geometry of the slide (Fig. 2). These data also allow the headwall and a lateral margin (Eastern Lateral Margin, Fig. 2) of the slide to be determined (not covered by the 3D seismic reflection data). The 3D seismic reflection data cover most of the toe domain of the slide (Fig. 2). Mapping of the seabed and basal shear surface of the slide enables us to constrain the structural style of its toe domain and infer the processes occurring during transport and cessation of the slide. The following seismic attributes were used to visualise the range of intra-MTC structures: (i) *variance* – to show discontinuities such as imbricated thrusts (e.g. Chopra & Marfurt 2007); and (ii) *spectral decomposition (RGB blending)* – to highlight heterogeneities of internal body of the slide by blending three bins of frequency volume with assigned colours (i.e. red, green and blue represent low, mid and high frequencies, respectively) (e.g. Partyka et al. 1999; Eckersley et al. 2018). We extracted these attributes along an iso-proportional slice (see Zeng et al. 1998) and horizontal time-slices, thereby generating map-view images of seismic facies and structural variability.

THE HAYA SLIDE

General characteristics

The Haya Slide is located *c.* 10 km off the coast of Sulawesi (Figs. 1 and 2). It is *c.* 16 km long, extending south-westwards from the lower slope (*c.* 1700 m below sea-level) to the basin floor (*c.* 2000 mbsl). The slide has a lobate geometry (Fig. 2): (i) it is *c.* 7 km wide in its headward region on the lower slope, (ii) widens to *c.* 15 km along its frontal margin in the centre of the basin floor, and (iii) it covers an area of 150 km^2. The slide was derived from the southern flank of a thrust-cored anticline within the SSP. The anticline has a broadly arcuate trend and is dissected by the headwall of the slide (i.e. extending

from 1700 to 1900 mbsl (Fig. 2a). The external limits of the slide are defined as follows (Fig. 2): (i) Northern Lateral Margin, (ii) Eastern Lateral Margin, and (iii) Frontal Margin. This external geometry, including the inferred headwall erosion on the southern flank of the seabed anticline, indicates that the slide was emplaced towards the SW. Correlation with the laterally equivalent, slide-hosting package in wells XR-1 and XS-1 (Fig. 1d), confirms that the slide is located within the Quaternary. Cores from the slope (TGS009) and basin floor (TGS194) locations (Fig. 1b) indicate that (i) slope sediments are composed of clayey (fine to medium) sand, with low-medium cohesion and medium-high water content, and (ii) basin floor sediments are characterised by very soft to firm clay, with medium cohesion and medium-high water content.

The 3D seismic reflection data cover *c.* 78% of the slide, mainly covering its downdip portion and excluding the headwall region (see inset map in Fig. 3a). Thickness patterns (Fig. 3a) and frequency characteristics (Fig. 3b) display gradual variations in both strike and dip directions, which enable subdivision of the slide. Strike-oriented thickness variations highlight three distinct areas (Fig. 3a): (i) A (*c.* 170-200 m thick), (ii) B (*c.* 140-170 m), and (iii) C (*c.* 70-140 m). All three areas thin and wedge-out abruptly downdip, at approximately the same rate, towards the Frontal Margin. Area C also thins rapidly, at a similar rate, towards the Northern Lateral Margin that represents a boundary separating the downslope-translating slide and stationary substrate. Note that the Eastern Lateral Margin is inferred using bathymetry data alone, whereas the Northern Lateral Margin is imaged directly by the 3D seismic reflection data.

Dip-oriented variations are defined by an isoproportional slice, taken midway between the basal shear surface and seabed (Fig. 3b), which shows frequency changes indicative of seismic facies and/or structural variability. The inner part of the slide is characterised by an overall lower RGB blend frequency and relatively short, discontinuous along-strike lineations. In contrast, outer areas display higher RGB blend frequency with longer, more continuous lineations, which extend across Areas A-C (Fig. 3b). These lineations predominantly trend E (090-270^{o}) in the south (Area A) and N to NW (000-180^{o}, 020-200^{o}) in the west (Area C).

Three dip-oriented seismic sections across Areas A, B and C, oriented perpendicular to the curved lineations (Fig. 3b), define the internal character of the slide (Fig. 4a-c). These sections show that the inner part of the slide comprises chaotic, highly discontinuous, low-amplitudes reflections, which corresponds to the low RGB blend frequency seen in the spectral decomposition map (Fig. 3b). This seismic expression is typical of an internally disorganised and highly deformed debrite, as compared to other, drilled examples of MTCs (e.g. Piper et al. 1997; Posamentier & Martinsen 2011). Between the inner and outer parts, the isolated high RGB blend frequency bodies on Figure 3b (which dominate the outer part) correlate with isolated, folded, high-amplitude reflections encased within the inferred debrite (Fig. 4a-c). These bodies are interpreted as megaclasts, with their long axes oriented sub-parallel to the curved lineations (Jackson 2011; Alves 2015; Gamboa & Alves 2015; Hodgson et al. 2018; Sobiesiak et al. 2018; Sobiesiak et al. 2019). The more continuous curved lineations in the outer part of the slide (Figure 3b) correspond to pairs of sharp discontinuities within the slide. These discontinuities converge downward onto the basal shear surface (e.g. Fig. 4a). We interpret these discontinuities as forethrusts (i.e. NE-dipping) and backthrusts (SW-dipping) that bound folded 'pop-up blocks' in their hangingwalls (e.g. Frey-Martínez et al. 2006; Bull & Cartwright 2019). The pop-up blocks (and their adjacent footwalls) are the lineation-bound, high RGB blend frequency bodies seen in map-view (Fig. 3b). Pop-up blocks are offset along-strike by discontinuities trending oblique (170°-350°) to the general trend of the bounding thrusts (Fig. 3b). These discontinuities are interpreted as sub-orthogonal shear zones (*sensu* Steventon et al. 2019) that may record boundaries between different flow-cells that moved at different speed within the translating failed mass (e.g. Masson et al. 1993; Steventon et al. 2019). They are not interpreted as longitudinal shears (*sensu* Bull et al. 2009b) that are sub-parallel to the transport direction.

Although thrust-bound pop-up blocks typify the outer part of the slide, there are significant lateral variations (from Area A to Area C) in structural style and seismic facies characteristics, which are described below.

Area A

Characteristics of Area A

A gradual downslope-deepening of the basal shear surface characterises the base of the slide in Area A. The surface steps up to form a steep ramp (*c.* 60°) that defines the slide's frontal margin (Fig. 4a). The basal shear surface is deepest (*c.* 200 mbsf) adjacent to the frontal margin, with the basal shear surface essentially being horizontal. The upper surface of the slide is of low relief in the inner part, which may partly reflect the filling of top-surface relief by post-emplacement sedimentation (ponded sediments in Fig. 4a). The upper surface becomes more rugose down-dip and reaches its highest relief (15 m) at the frontal margin.

Seismic reflections in the outer part of the slide in Area A are well-imaged and can be directly correlated with undeformed strata beyond the frontal margin, despite being contractionally offset by thrust faults (Fig. 4a). The internal reflections of the slide become more irregular, and harder to trace, towards the inner part. In area A, the average throw and dip of the fore- and backthrusts are *c.* 30 m and *c.* 45°, respectively, with the spacing between thrust pairs (measured from crest to crest of pop-up blocks) ranging from 400 to 500 m.

Interpretation of Area A

The steep frontal ramp that separates undeformed basin-floor strata from the slide is a classic frontally-confined (*sensu* Frey-Martínez et al. 2006) termination style (Fig. 4a). Similar to previously documented frontally-confined MTCs, the thickness of the slide in the outer area (*c.* 200 m) is only expressed by minimal seabed relief at the edge of the deposit (*c.* 15 m) (e.g. Lastras et al. 2004; Frey-Martinez et al. 2005).

Internal reflections show higher preservation of stratal reflections in the outer than the inner parts, suggesting that the youngest thrust is located at the frontal margin of the slide (Fig. 4a), similar to those observed from outcrops (e.g. Alsop et al. 2019) and seismic reflection data (e.g. Frey-Martínez et al. 2006; Bull & Cartwright 2019). Physical modelling results suggest that regular spacing of fore-

and backthrusts is indicative of an MTC that was translated on a low friction basal shear surface (Huiqi et al. 1992).

Area B

Characteristics of Area B

Gradual downslope-deepening of the basal shear surface is also observed in Area B, with the surface progressively stepping up through stratigraphy to define three stepped levels of frontal ramp (Fig. 3a and Fig. 4b). The basal shear surface is deepest (*c.* 170 mbsf) immediately upslope from the first and deepest frontal ramp with the highest relief (30 m). The other two ramps are more gently-dipping and have lower relief (*c.* 20 m) (Fig. 4b). These three ramps truncate otherwise continuous, sub-parallel reflections defining the pre-slide substrate. The substrate in Area B dips very gently (*c.* 1^{o}) in an opposing direction (i.e. north-eastwards) to the slide transport direction. The seabed in Area B is smooth but becomes more rugose downdip (Fig. 4b). Most notably, the highest seabed relief (*c.* 10 m) is located immediately above the deepest point of the basal shear surface.

The nature and distribution of the seismic facies in Area B differs from those of Area A, most notably a much higher level of reflection discontinuity. Also, the least disturbed strata (i.e. semi-continuous seismic reflections) occur in the central part of the slide, immediately upslope from the first frontal ramp. Directly above the frontal ramps, reflections are extremely chaotic with variable, higher amplitude seismic facies encased within more extensive transparent seismic intervals, which resemble those in the inner part (Fig. 4b). In the central area, where stratal reflections have the highest preservation, pop-up blocks and thrusts are geometrically similar to those in Area A. However, these pop-up blocks have a spacing of *c.* 150-300 m, which is about half that of Area A. Measuring the throw and dip of thrusts in Area B is harder than in Area A, likely due to seismic resolution limitations and the closer spacing of the thrusts. Where we can trace a marker horizon between thrust-bound pop-ups, the throw and dip of the thrusts are 49 m and 60^{o}, respectively (i.e. similar to the maximum values observed in Area A).

A distinctive upstanding, undeformed block is identified on a variance timeslice and seismic section (see 'Intact block' in Fig. 5), which marks the transition between Area A and B. This block extends gradationally downwards into the undeformed slope-to-basinfloor strata (Fig. 5b), which continue unbroken towards the east. The block is bound (north and south) by the steep frontal ramp defining Area A and pop-up blocks within the toe domain of the slide (west and south). The block is capped by sub-parallel, variable-amplitude reflections, while further south it is bound by folded reflections that are cross-cut by minor thrusts. These thrusts detach onto a reflection that is stratigraphically shallower than the basal shear surface (Fig. 5b). We suggest that the basal shear surface steps up above this feature, before stepping down onto the reflection onto which the minor thrusts detach. The surface then steps up again to define the outermost frontal margin in Area B. Beyond this outermost frontal margin, a gently folded reflection is observed that probably marks the position where the next thrust would have formed (Frey-Martínez et al. 2006).

Interpretation of Area B

The stepped geometry of the basal shear surface confining the slide in Area B argues against frontal emergence of the slide (Frey-Martínez et al. 2006). Seismic facies above the stepped frontal ramp comprise variable-amplitude, somewhat chaotic reflections that resemble debrites (*cf.* Posamentier & Kolla 2003; Ortiz-Karpf et al. 2017) (Fig. 4b). Pop-up blocks in Area B are located immediately updip from the frontal ramps (Fig. 4b). Here, the slide is thinner, and it contains more closely-spaced pop-up blocks than those in Area A. We therefore speculate that there might be a relationship between thickness and pop-up block width/thrust fault spacing. This is consistent with the physical and numerical modelling by Liu & Dixon (1995), who demonstrate a positive linear relationship between thrust spacing and thickness of the strata.

We interpret the intact block between Areas A and B as a piece of *in situ* substrate, based on its lack of deformation and gradational seismic facies relationship with underlying and adjacent basin floor strata. Hence, it can be interpreted as a remnant block (*sensu* Bull et al. 2009b). Minor thrusts downdip

from the remnant block suggest that there is a zone of relatively high strain beyond the main body of the slide (Fig. 5b). This zone of high strain could be a distributed shear zone, where compressional stress is transmitted beyond the frontal ramp (Hodgson et al. 2018). However, in those cases, the distributed shear zone is commonly in direct contact with the frontal margin of the main body (e.g. Watt et al. 2012). In our case, the remnant block exists in between two zones of relatively high strain (Fig. 5b). Therefore, an alternative interpretation is that the minor thrusts represent the lateral propagation of thrusts eastwards from Area C (Fig. 5a). This interpretation is plausible given that minor thrusts can be traced westwards on the variance time-slice, towards the main body of the slide (i.e. into Area C, Fig. 5a). The relationship between the main body of the slide, the remnant block, and the minor thrusts, partially resemble a process referred to as 'enveloping' (Hodgson et al. 2018). For example, a remnant block could form when an uneven frontal margin to the slide envelopes a large piece of substrate, but with the process terminating prior to complete entrainment of the block due to cessation of the slide's translation.

Area C

Characteristics of Area C

The basal shear surface in the outer part of Area C exhibits a similar geometry and internal characteristics to that of Area B, especially the staircase-like geometry of the basal shear surface (Fig. 4c). However, the basal shear surface here is associated with a pronounced change in dip and dip direction, defined by a change from *c.* 1° basinward dip to a *c.* 3° landward dip (Figs. 4c and 6a). This change in dip coincides with the deepest (120 mbsf) occurrence of the basal shear surface. The seabed in Area C is characterised by a (i) *c.* 10 m vertical relief, and (ii) a *c.* 6 km long and 2 km wide 'bulge', immediately updip of the slide's frontal margin (Figs. 4c, 6b-c). Adjacent to the Northern Lateral Margin, the basal shear surface is relatively flat, and the seabed shows rugosity similar to that in Areas A and B, but with a shorter wavelength (Fig. 6d).

The internal characteristics of the slide in Area C, which resemble those in Area B, comprise the following: (i) chaotic reflections of variable amplitude encased within very low-amplitude reflections at the frontal margin, (ii) pop-up blocks within the slide's outer part, and (iii) megaclast-bearing debrites in the inner part (Fig. 4c). However, the pop-up blocks in Area C are more closely spaced (*c.* 100-150 m) than those in Area B, which results in low stratal preservation in seismic sections (Fig. 4c). Thus, despite being well-imaged in map-view, from which pop-up blocks spacing can be measured (Fig. 3b), dip and throw measurements in Area C are uncertain (Fig. 4c). The frontal margin in Area C is characterised by rapid pinch-out of the slide's internal body onto the inclined (*c.* 3^{o}) substrate (Fig. 4c). Towards the Northern Lateral Margin, the spacing between pop-up blocks is even shorter (*c.* 70-100 m), and the basal shear surface is shallower (70 mbsf) (Figs. 3 and 6d). Near the frontal margin, sub-parallel, discontinuous, high-amplitude reflections occur between the basal shear surface and the largely transparent seismic facies defining the main body of the slide (Fig. 4c). These reflections are correlated with a *c.* 25 m-thick interval located basinward of the slide, comprising inclined, largely undeformed, reflections (Fig. 4c).

The boundary between Areas B and C comprises a NE-trending/NW-facing ramp, which is laterally continuous with the NW-trending/NE-facing frontal ramp of Area B (Fig. 7a). Variance attributes extracted from a 50 ms TWT thick window above the basal shear surface show several NW-trending lineations that terminate against the NE-trending ramp. In seismic section, these lineations correspond to fold-and-thrust belt structures in Area C (Fig. 7b). Thus, the NE-trending ramp forms a boundary between the fold-and-thrust system and the undeformed substrate. The NE-trending ramp also coincides with a positive relief on the seabed.

Interpretation of Area C

The slope gradient break at the basal shear surface and emergent of the leading-edge part of the slide that onlaps onto the underlying inclined substrate are likely to be related. We suggest that the physical impact of the downslope-translating slide onto its substrate was highest where the basal shear surface

abruptly changes dip and dip direction (Ogata et al. 2014b). Following this impact, variations in the mechanical properties of the substrate likely controlled the morphology of the basal shear surface (Strachan 2002; Frey-Martinez et al. 2005; Moernaut & De Batist 2011). For instance, substrates with higher shear strengths (e.g. due to lower pore-pressure) force the basal shear surface to step-up to shallower substrates and propagate along inclined substrates that have lower shear strength (Fig. 4c). The inclined basal shear surface and momentum gained by the slide at the dip change provide sufficient inertial energy for the translating mass to abandon the basal shear surface and emerge onto the coeval basin floor, and to onlap the bathymetric high (Figs. 4c, 6b) (Frey-Martinez et al. 2005; Frey-Martínez et al. 2006). Therefore, we classify the slide in Area C as frontally-emergent (*sensu* Frey-Martinez et al. 2006). However, the slide also becomes frontally-confined adjacent to the Northern Lateral Margin, where the slide is thin, and the basal shear surface is relatively flat and lacks a distinct dip change (Fig. 6d; *cf.* Area A in Fig. 4a).

The abrupt change in basal shear surface dip has at least two additional consequences. Firstly, the internal body of the slide was likely disaggregated due to the buttressing effect of the underlying substrate (Mandl & Crans 1981). This resulted in the partially-disaggregated debrite facies in the frontal margin area, which is manifested as the broad bulge on the seabed (Fig. 6b-c). Secondly, the impact of the translating mass onto the substrate develops a zone of stratigraphically parallel, discontinuous reflections directly on top of the basal shear surface (e.g. Joanne et al. 2013). We interpret these reflections as lying within the basal shear zone, in which the substrate was deformed due to compressional forces exerted by the slide, but was not fully entrained (e.g. Hodgson et al. 2018; Cardona et al. 2020).

The abrupt boundary between Areas B and C indicates that the basal shear surface evolved differently between the two areas, where the frontal ramp of Area B was cross-cut by the main body in Area C (Fig. 7a). This cross-cutting relationship probably formed by the slide's cannibalisation of the substrate in Area C, which formed the NW-facing ramp (Fig. 7a-b). Lateral variations in basal shear surface

growth and geometry could also be related to lateral variations in the mechanical properties of the stratigraphy overlying the basal shear surface (e.g. permeability, pore-pressure and related shear strength). In addition, variations in the magnitude of stress exerted by the slide onto, and into, the substrate in adjacent areas may have occurred (Strachan 2002; Frey-Martinez et al. 2005). Positive seabed relief adjacent to the NE-trending ramp likely reflects a buttressing effect of the main body of the slide against the ramp as new material was entrained by the slide (Fig. 7b).

Estimation of translation distance and along-strike variability of strain in the toe domain

We here estimate the translation distance of the Haya Slide based on an assessment of shortening within Area A that has the best preservation of internal reflections. We also quantify intra-MTC strain of a pop-up block within Area A to investigate how strain varies along strike.

Shortening strain and its relationship to translation distance

The distance travelled by the slide where frontally confined can be estimated by measuring total shortening in the toe domain, as long as the fold-and-thrust belts and the internal reflections are well-preserved and imaged (*cf.* Frey-Martínez et al. 2006; Bull & Cartwright 2019). We note that the calculated translation distance here is a first-degree estimation of how far the slide has travelled in the toe domain (Frey-Martínez et al. 2006), and, thus, it does not represent run-out distance, which is measured from the headwall to the leading-edge of the deposit (Clare et al. 2018). We note that the estimated shortening values would underestimate intra-MTC strain (i.e. minimum value), as the line-length method does not account for shortening within pop-up blocks due to sub-seismic strain, lateral compaction accommodated by porosity loss via dewatering, and/or grain crushing (Moore et al. 2011; Armandita et al. 2015; Alsop et al. 2019; Steventon et al. 2019).

A representative depth-converted seismic-section in Area A (interval velocity derived from wells XR-1 and XS-1) was selected for our shortening calculation based line-length method (see Figs. 3b and 4a). This section is orientated perpendicular to the strike of the fold-and-thrust belt, and stratal reflections within individual thrust-bound blocks are well-imaged, and can thus be interpreted with confidence.

Two intra-MTC horizons were interpreted (H1-2, see Fig. 4a) to better constrain the amount of horizontal shortening and to determine how this varies vertically. These horizons extend from undeformed basin-floor strata to the updip limit of the outer part (Fig. 4a). The present and restored lengths of H1, the deepest horizon, are 6.73 km and 7.79 km, respectively, which equate to 14% contraction (1.06 km). In contrast, the shallower H2 horizon experienced only 8% contraction (0.61 km), derived from present and initial lengths of 6.65 km and 7.26 km, respectively. This analysis shows two key results: (i) contractional structures in Area A (Fig. 4a) formed in response to horizontal translation of the slide over a relatively short distance (0.61-1.06 km), and (ii) greater contraction of the deeper H1 horizon compared to the shallower H2 indicates depth-dependent layer shortening due to penetrative strain (e.g. thickening and partial area loss of deeper layer, Koyi 1995).

Along-strike variability of intra-MTC strain

An along-strike analysis enables the kinematics behind the spatial configuration of fold-and-thrust belts to be assessed (Dahlstrom 1969). Such studies have been performed for kilometre-scale, deep-water fold-and-thrust belts using 3D seismic reflection data (e.g. Higgins et al. 2009; Totake et al. 2018). Here, we document the along-strike variability of intra-MTC strain at a significantly smaller-scale, but exceptionally well-imaged, fold-thrust system within the Haya Slide.

We conducted the along-strike analysis on Pop-up Block 3 (i.e. the third block counted from the frontal margin, and herein referred to as PB-3; see Fig. 4a) and its associated fore- and backthrusts. This pop-up block is ideal for this analysis because its main bounding thrust fault (FT-1) and Horizon H2 can be interpreted over the longest distance (*c.* 3 km along strike, see Fig. 8a); other pop-up blocks are shorter and more segmented along strike (*c.* 0.5-1 km). We measured throw along the fore- (FT) and backthrusts (BT) at intervals of 20-200 m. The throw was measured because most thrusts dip 40^{o}-60^{o} and strain is better quantified by measuring throw, rather than heave that diminishes with increasing dip (Totake et al. 2018).

374 Mapping of H2 laterally from the representative section of Area A (i.e. Fig. 4a) reveals a more
375 complicated configuration of pop-up structures associated with PB-3; whereas there is only a single
376 pop-up in the east (PB-3a), there are two in the west (PB-3b-c; Fig. 8a). These three pop-up blocks are
377 readily identified on a variance time-slice (Fig. 8b). Here, one of the sub-orthogonal shear zones
378 identified in the previous section (see General Characteristics and white dotted lines in Fig. 3b), trends
379 oblique to, and cross-cuts, the thrust faults near the central part of the focused study area (white
380 dotted line in Fig. 8b). This shear zone clearly defines the boundary between PB-3a in the east and PB-
381 3b and c in the west (Fig. 8a).

382 The structural configuration is seen on both the H2 time-structure map and variance time-slice (Fig.
383 8a-b). At the shear zone, the southern margin of the PB-3a and b shows an 80 m left-lateral (sinistral)
384 offset (Fig. 8b). Therefore, the translated mass to the west appeared to travel downdip only a small
385 amount further than the mass to the east when compared to the overall estimated translation
386 distance of the slide (i.e. 7.5-13% of 0.61-1.06 km translation distance). PB-3a is bound on its northern
387 margin by one major backthrust (BT-1), and one minor FT-2 exists adjacent to FT-1. In contrast, PB-3b
388 is bound on its northern side by BT-2 and -3 that forms a 'soft-linkage' with each other (*sensu* Walsh
389 & Watterson 1991). Unlike PB-3a and -b, PB-3c is not bound by FT-1, but is instead bound by two
390 forethrusts (FT-4 and FT-5) and two backthrusts (BT-4 and BT-5). BT-1 and BT-4 are soft-linked (near
391 the shear zone) and bound limit the northern margin of PB-3a and c, respectively, (Fig. 8a). The faults
392 bounding the three pop-up structures generally strike E-W to ESE-WNW. In addition to the faults that
393 define PB-3a-c, we identify two faults (i.e. FT-3 and BT-6) within the shear zone (Fig. 8a-b). These faults
394 bound a narrow (*c.* 100 m-wide), uplifted block that may have formed due to transpression within the
395 shear zone (Sanderson & Marchini 1984).

396 Throw profiles of individual fore- and backthrust faults show multiple maxima and minima (Fig. 8c),
397 resembling larger, tectonic-scale fold-thrust systems, such as in offshore NW Borneo (Totake et al.
398 2018) and Niger Delta (Higgins et al. 2009). T-x plot of FT-1 shows that it has a slightly lower throw (*c.*

5-10 m) in the western (PB-3b) than in the eastern (PB-3a) domains (Fig. 8c). This contrasts with an increase of the number of thrusts, resulting in a significantly higher cumulative throw: from *c.* 20-40 m in the E to *c.* 40-60 m in the W (Fig. 8c). This might indicate that pop-up structures in the western domain are in a more advanced phase of growth (e.g. Cartwright et al. 1995; Totake et al. 2018). A local minima in the cumulative throw profile, which coincides with the local minima of FT-3, marks the boundary between the eastern and western domains (Fig. 8c). This boundary may represent a paleo-linkage site (Ellis & Dunlap 1988), which in this study coincides with the shear zone (Fig. 8a-b). The seismic sections across PB-3 depict the change in the fold-and-thrust configuration along strike (Fig. 8d-f), from the eastern area, the shear zone, to the western area.

We found two different strain domains within the translated mass: in the eastern and western domains, separated by an intra-MTC, syn-emplacement shear zone (Fig. 8a-b). These two domains were likely transported a similar distance. However, they experienced significantly different amounts of contraction, as evidenced by the cumulative throw profile (Fig. 8c). Several studies show that internal shear zone reflects differential timing or velocities of translating masses (or 'cells') within an MTC (Masson et al. 1993; Gee et al. 2005; Bull et al. 2009a; Steventon et al. 2019). It implies that a cell that is transported for a longer duration, or at a higher velocity, would likely travel further. We show here that when the intra-MTC cells could not travel further due to frontal confinement of the failed mass, stress imposed by still-moving material towards its rear can be accommodated by the formation of additional/larger contractional structures (e.g. folds and thrust faults). This process results in along-strike variability in the style and magnitude of intra-MTC strain, with the shear zone separating the intra-MTC cells recording the different amounts of strain.

DISCUSSION

We here discuss the slide transport processes and lateral variability of frontal emplacement and intra-MTC strain within the toe domain. Also, we discuss the implications for assessing the seal potential of MTCs in relation to hydrocarbon accumulations.

Modes of transport

Frey-Martínez et al. (2006) show the headwall domain of frontally-confined MTCs are defined by internally coherent, normal fault-bound blocks. In this domain, there is only limited depletion of the failed mass immediately downdip of the headwall. However, more recent studies show that major sediment depletion in the headwall domain can occur even if the MTCs are frontally confined (e.g. Lastras et al. 2004; Watt et al. 2012; Joanne et al. 2013). In such cases, these frontally-confined MTCs are generally characterised by strongly disaggregated, debritic material in their inner parts, rather than fault-bound blocks. Downdip, contractional structures (e.g. folds and imbricated thrusts) display increasing stratal preservation distally.

The Haya Slide comprises an inner, debrite-dominated part and an outer part dominated by contractional structures. The debrite likely originated from the collapse of the southern flank of an updip anticline (see Fig. 3). This deformed the seabed and entrained the substrate (Fig. 9a), which resulted in flow bulking further downslope (Gee et al. 2001; Butler & McCaffrey 2010). Substrate entrainment and subsequent downslope translation then produced transparent seismic facies (i.e. the debrite in Fig. 4), indicating that the incorporated material was increasingly disaggregated (Posamentier & Kolla 2003; Ortiz-Karpf et al. 2017). Erosion and disaggregation by the debris flow continued until the shear stress exerted by the flow was unable to entrain more substrate (Fig. 9b). At this point, the debris flow applied significant shear and compressional stress (lateral loading) to the substrate ahead of, and to the sides of, the flow (Butler & McCaffrey 2010; Hodgson et al. 2018).

The strata ahead of the debris flow were translated a short distance (i.e. 0.61-1.06 km), forming broadly symmetrical pairs of fore- and backthrusts (Fig. 9c). This symmetrical geometry of the thrusts is likely due to low basal friction during shearing (Huiqi et al. 1992). The low basal friction may reflect the fact that the failed mass had a high water content, and thus high pore pressure (e.g. Armandita et al. 2015). The two styles of MTC-substrate interactions, i.e. erosion and deformation (Fig. 9c), have been documented elsewhere, both in seismic reflection (e.g. Schnellmann et al. 2005; Watt et al. 2012;

449 Joanne et al. 2013; Ogata et al. 2014a; Bull & Cartwright 2019; Omeru & Cartwright 2019; Steventon
450 et al. 2019), and field data (e.g. Van Der Merwe et al. 2011; Sobiesiak et al. 2016; Sobiesiak et al. 2019;
451 Cardona et al. 2020). Adjacent to the toewall, the basal shear surface evolves differently along strike
452 (Fig. 10), which will be discussed in the following section.

453 **Lateral variability of the toe domain**

454 *Lateral variability of frontal confinement*

455 Moernaut & De Batist (2011) investigated sub-lacustrine MTCs to understand what controls whether
456 an MTC remains confined, or whether it abandons its basal shear surface and emerges onto the coeval
457 basin floor. They conclude that the drop height and depth of the basal shear surface are the main
458 factors controlling frontal emplacement style. The former represents a driving force (i.e. gravitational
459 potential energy), and the latter represents a resisting force (i.e. potential energy needed to be
460 exceeded for the MTC to emerge).

461 The Haya Slide originated from a headwall at a depth of *c.* 1700 mbsl, and its frontal margin is at *c.*
462 2000 mbsl (the basinward extent of Areas A to C) (see Fig. 3). Thus, the drop height of the slide is 300
463 m, which provided a similar driving force (potential energy) for all the three frontal areas. However,
464 the depth of the basal shear surface, and thus the thickness of the slide, varies laterally: it is deepest
465 in Area A (*c.* 200 mbsf) and shallowest in Area C (*c.* 120 mbsf). This lateral variability of basal shear
466 surface depth, slide thickness and degree of confinement must also reflect lateral changes in the ratio
467 between the resisting and driving forces (Fig. 10). In particular, the driving forces needed for the slide's
468 emergence in Area A were greater than that in Area C. Therefore, the Haya Slide exhibits a lateral
469 variation of frontal emplacement (Fig. 10); i.e. full frontal confinement in Area A, partial confinement
470 across several staircase-like frontal ramps in Area B, to frontal emergence in Area C. Lateral friction
471 along the Northern Lateral Margin may have also locally increased the resisting force in addition to
472 the basal friction (e.g. Joanne et al. 2013), such that the slide is frontally-confined in that area despite
473 being at its thinnest (Fig. 6d).

There is also a broad correlation between the basal shear surface morphology (i.e. depth and slope gradient break) and the overlying structural style in the toe domain. In Area A, for example, a relatively flat gradient, coupled with a deep basal shear surface, is associated with a steep (*c.* 60°) frontal margin (Figs. 4a and 10). This steep frontal margin represents the youngest forethrust that was formed as the slide ceased to translate (Fig. 11a) (e.g. Watt et al. 2012; Joanne et al. 2013; Alsop et al. 2019). In contrast, Area C displays a low-angle (3°), upslope-dipping, and relatively shallow basal shear surface related to the frontal ramp and slide emergence onto the coeval basin floor (Figs. 4c and 10). Here, a bathymetric high (see Fig. 6a-c) that existed prior to slide emplacement formed inclined strata ahead of the slide. This inclination increased the impact of the slide onto the substrate. The increased impact led to: (i) the formation of basal shear zone, and (ii) allowed the slide to transfer remaining exerted stress by abandoning the basal shear surface and translate on the coeval seafloor (Fig. 11b). Such distal bathymetric confinement has also been documented elsewhere, for instance, in offshore Colombia, where channel-levee morphology could deflect and/or block debris flows (Ortiz-Karpf et al. 2017).

Areas A and C represent end-member styles of the sub-slide basal shear surfaces. Morphologically, the basal shear surface in Area B lies between Areas A and C, being defined by a low-angle (1°) surface, an intermediate-depth and a staircase-like set of frontal ramps (Fig. 4b and 10). The formation of these ramps can be compared to the ramps and flats present along non-planar thrust faults, where the ramps tend to form in relatively high-shear strength layers, and the flats (e.g. basal shear surface connecting the ramps) in weaker layers (Fossen 2016). The potential energy of the slide in Area B might have been progressively (rather than instantaneously) dissipated in the distal area (Fig. 11c). Here, the basal shear surface may have propagated downslope along a horizon until it encountered a layer with higher shear strength (i.e. the red point in Fig. 11c). At that point, the basal shear surface stepped-up through stratigraphy and continued to propagate in shallower levels (i.e. initiated from the green point in Fig. 11c). This process might have continued several times to form the staircase-like frontal ramps, eventually terminating when the shear strength of the strata ahead of the flow exceeded the shear stress exerted by the slide (Fig. 11c). Alternatively, the staircase-like geometry

might represent a transitional style between full frontal confinement and full frontal emergence. The first frontal ramp in Area B links along-strike to the frontal ramp in Area A (Fig. 3a). Thus, this first step can be interpreted as the initial toewall. However, this initial toewall was not developed to form a steep ramp such as that in Area A. Instead, the debrite-like seismic facies above the subsequent steps might represent a style of frontal emergence (Fig. 4b). Consequently, the slide must have abandoned the basal shear surface, and progressively shallowed and incorporated material downdip from the initial toewall. This differs to Area C where the slide expelled material on to the coeval basin floor.

There is also some degree of correlation between the depth of the basal shear surface and the degree of disaggregation adjacent to the toewall. In Area A, where the basal shear surface is deeply-rooted, internal reflections of the slide are well-preserved (Fig. 11a). In contrast, in Areas B and C, where the basal shear surface progressively shallows, internal reflections of the slide exhibit debritic facies, indicating internal disaggregation (Fig. 11b-c). A similar relationship has also been documented in the thinner part of MTCs in offshore Brazil (Alves & Cartwright 2009; Gamboa et al. 2011) and offshore Colombia (Ortiz-Karpf et al. 2017). These studies conclude that the shallowing basal shear surface led to an increase in shear stress at the base of the flow with increased disaggregation.

Hence, we conclude that the interplay between stresses exerted by parent flow and variation of mechanical properties of the sedimentary package (both locally and regionally), controls the morphology of the basal shear surface (Figs. 10 and 11) (Bull et al. 2009b; Shanmugam 2015; Hodgson et al. 2018; Sobiesiak et al. 2018).

Lateral variability of intra-MTC strain

Only a few studies have used seismic reflection data to quantify intra-MTC strain (Bull & Cartwright 2019; Steventon et al. 2019). More specifically, these studies have focused on: (i) strain balancing between headwall and toe domains of MTC located offshore Uruguay (Steventon et al. 2019) and offshore Norway (i.e. Confined Stroregga Slide (CSS), Bull & Cartwright 2019); and (ii) assessment of depth-dependant layer shortening in the toe domain (Steventon et al. 2019). The Uruguay example

shows that contractional strain in the toe domain is apparently greater than (by *c.* 3-14%), and thus does not balance, extensional strain in the headwall domain (Steventon et al. 2019). This strain deficit is attributed to sub-seismic penetrative strain, likely associated with grain-scale deformation, and porosity and fluid loss (Koyi 1995; Koyi et al. 2004; Burberry 2015; Dalton et al. 2017; Alsop et al. 2019). In contrast, the study of the CSS found that extensive sediment depletion in the headwall domain is accommodated by only relatively mild contraction (*c.* 5%) in the toe domain (Bull & Cartwright 2019). This discrepancy is inferred to reflect a subsequent phase of deformation that involved the removal of a significant amount of material from the headwall domain after emplacement of the CSS.

Besides longitudinal balancing of MTCs, seismic-scale vertical variability of intra-MTC strain has also been documented. Steventon et al. (2019) documented that the deeper horizon (i.e. closer to the basal shear surface) experienced more shortening (*c.* 27%) than the shallower horizons (*c.* 18%) in the toe domain of the MTC, offshore Uruguay. We find similar results in the Haya Slide, where deeper (H1) and shallower (H2) horizons record *c.* 14% and *c.* 8% of shortening, respectively (Fig. 4a). These observations suggest that the magnitude of shortening estimate depends on the measurement depth due to depth-dependant horizontal shortening, with strain being greatest at depth. Physical models of horizontal shortening suggest that the increase of shortening with depth is balanced by bed-length decrease, lateral compaction of deeper layers, layer-normal thickening of shallower layers, and increased thrust displacement (Koyi 1995; Koyi et al. 2004; Burberry 2015). One or a combination of these processes might occur within the toe domain of a seismic-scale MTC.

The examples above show that intra-MTC strain varies both longitudinally and vertically. Our along-strike analysis of PB-3 and its associated thrusts indicate that intra-MTC strain also varies laterally, with a shear zone separating two regimes of contraction within a translated mass (Fig. 8). This represents a seismic-scale example of the field data-derived, multi-cell flow model of Alsop & Marco (2014) (see also Farrell 1984). This model states that a first-order, single-cell MTC is composed of many smaller, second-order flow cells that are formed during translation and may locally interact (Alsop &

550 Marco 2014). This local interaction is revealed by our along-strike analysis of PB-3, which we infer is
551 contained within a more extensive, first-order cell. The eastern and western domains of the pop-up
552 block represent second-order flow cells, with the shear zone representing the flow cell boundary. In
553 the context of multi-cell flow model, PB-3 might initially have been a single body (or cell) of sediment,
554 experiencing the same amount of stress laterally. However, velocity perturbations during translation
555 of the first-order cell, perhaps due to variable basal shear stress, initiated the formation of the shear
556 zone and caused formation of the two second-order flow cells within the initially continuous pop-up
557 block. The western cell may have been translated at a faster speed than the eastern cell, so that the
558 western cell records a more advanced stage of contraction than the eastern one.

559 **Impact of intra-MTC strain on seal potential**

560 MTCs can play at least two roles in the development of petroleum systems: (i) commonly as seals
561 (Algar et al. 2011, Cardona et al. 2016), and (ii) rarely as reservoirs (Sawyer et al. 2007; Algar et al.
562 2011; Shanmugam 2012; Arfai et al. 2016; Cardona et al. 2016). This is controlled by three key
563 parameters: (i) provenance lithology, most notably sand/mud ratio (Jenner et al. 2007; Omosanya &
564 Alves 2013), (ii) substrate lithology and erodibility (Cardona et al. 2020), and (iii) the degree of internal
565 disaggregation (Alves et al. 2014), which may include significant permeability reduction (2.5-3.5 orders
566 of magnitude) due to grain crushing in otherwise good quality reservoirs (Crawford 1998). A strongly
567 disaggregated (e.g. debrite-dominated) MTC derived from a very-fined grained source may result in
568 high seal potential (Alves et al. 2014; Omeru 2014; Cardona et al. 2016). However, this may be
569 compromised by entrainment of coarser-grained substrate and/or by the presence of rafted blocks
570 (megaclasts) with reservoir potential (Gamboa & Alves 2015; Cardona et al. 2016; Cardona et al. 2020).
571 Hence, substrate entrainment may result in (i) increased permeability and thus lower overall seal
572 capacity, and/or (ii) localized high-permeability zones that can promote vertical fluid migration and
573 hydrocarbon leakage.

Seal competence can vary longitudinally, from head to toe domains of the MTC, due to substrate entrainment and shearing during transport (e.g. Cardona et al. 2020). The Haya Slide is a clay-rich MTC that contains debritic facies in the inner part; this area may therefore represent a good hydrocarbon seal when compared to the imbricated, but otherwise internally moderately undeformed blocks present in the outer part (Figs. 3b and 4). However, we also document notable along-strike variations in seismic facies in the outer part (Fig. 4). For instance, Area A is characterised by imbricated thrusts. If these thrusts lack clay smear and are relatively permeable compared to the flanking, very fine-grained host rock, they may be conduits for fluid migration, implying a higher seal risk for this area (i.e. low seal potential). Towards Area C, seismic facies become more chaotic and transparent, suggesting a higher degree of deformation and internal disaggregation. Seismic facies in Area C may thus suggest a better seal potential here than in Area A because chaotic and transparent seismic facies have higher seal potential than blocky MTCs containing preserved stratigraphy (Alves et al. 2014; Omeru 2014). Therefore, our results suggest that seal potential of an MTC can vary along both depositional dip and strike within any one domain. The results presented here could be used to inform prediction of seal potential in MTCs at deeper, economic burial depths.

CONCLUSIONS

A recent mass-transport complex (MTC), the Haya Slide, has been characterised in the Makassar Strait based on high-quality 3D seismic reflection and bathymetry data. The slide originated from the collapsed flank of an anticline in the NE and transported radially to the SW. An along-strike analysis of the toe domain of the slide has provided the following conclusions concerning lateral variability of frontal emplacement and intra-MTC strain distribution:

1. The inner part of the toe domain is characterised by a debrite, which passes, first, downdip into megaclast-bearing debrite and, second, into coherent pop-up blocks towards the outer part. The debrite and the pop-up blocks are genetically-related, sharing the same bounding surfaces. Lateral loading by the debrite onto coherent strata induced progressive downslope

599 failure. Shortening estimates across the coherent strata show 8-14% of shortening, equating
600 to 0.6-1.1 km of downslope translation.

601 2. The outer part of the toe domain exhibits the following lateral variability: (i) depth and
602 gradient of the basal shear surface, (ii) trend and spacing of the pop-up blocks and their
603 associated thrust faults, and (iii) frontal emplacement processes. A deep and relatively flat
604 basal shear surface is associated with frontal confinement, where steep ramp separates
605 undeformed strata and the slide. A shallow and upflow-dipping basal shear surface is
606 associated with frontal emergence of the slide onto the coeval basin floor. Between these two
607 extremes, the frontal geometry is characterised by staircase-like frontal ramps. Internal
608 architecture of the slide may also be related to the geometry of the basal shear surface, where
609 highly disaggregated material can be associated with the progressive shallowing basal shear
610 surface. The interplay between drop height (i.e. driving force),= and along-strike depth
611 variation of basal shear surface (i.e. resistive force),= likely to determine the lateral variability
612 of frontal geometry of the slide. For instance, where resistive force < driving force led to
613 frontal emergence, otherwise the slide would be frontally confined.

614 3. A detailed study of fold-and-thrust structures within the region of pop-up block shows along-
615 strike variability of intra-MTC strain. This shows western and eastern regions of the toe
616 domain, separated by a shear zone, experiencing different amounts of contraction. The
617 western regime records a higher amount of strain, reflecting a more advanced phase of
618 structural growth compared to its eastern counterpart.

619 4. Previous studies have shown that an MTC could have variable seal competence based on its
620 axial domains (headwall to toe) due to different degree of disaggregation and substrate
621 entrainment. The Haya Slide shows not only that, but there are also distinctive along-strike
622 variations, which could influence seal potential predictions in petroleum systems.

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CONFLICT OF INTEREST

No conflict of interest declared.

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933 **FIGURE CAPTIONS**

934 **Fig. 1.** Geological setting and location map of the study area. **(a)** The Makassar Strait is surrounded by
935 tectonically active regions, where Eurasia, Indo-Australia, Philippine Sea and Pacific plates interact. A
936 strong ocean current flowing from Pacific towards Indian oceans, Indonesia Throughflow (ITF), flows
937 through the Makassar Strait (red arrow). **(b)** The study area is located in the southern end of Labani
938 Channel, a narrow strait between Paternoster Platform and western Sulawesi margin. The channel
939 connects the North and South Makassar basins. Major structural features adjacent to the study area
940 are fault zones (Palu-Koro and Paternoster fault zones) and fold-thrust belts (e.g. Brackenridge et al.,
941 2020; Cloke et al., 1999). The fold-thrust belts are divided into the Northern (NSP), Central (CSP) and
942 Southern (SSP) structural provinces (Puspita et al., 2005). The dark blue line marks the extent of 3D
943 seismic reflection data, and the green line outlines the area covered by multibeam data. Two green
944 dots represent wells within the seismic reflection data. The small, yellow area marks the extent of the
945 Haya Slide (see Fig. 2). Blue and red dots are the location of near-seabed sediment cores of TGS009
946 and TGS194, respectively. **(c)** A conceptual cross-section across the Makassar Strait showing MTCs
947 accumulation in the basin and their related sources, i.e. accretionary shelf (related to Mahakam Delta)
948 in the West and collapse of anticline flanks in the East. Inferred based on Puspita et al. (2005) and
949 Brackenridge et al. (2020). **(d)** A seismic line correlating the Haya Slide (yellow-shaded) and the two
950 wells (i.e. XS-1 and XR-1). The sedimentary package containing the slide is Quaternary in age and
951 predominantly consists of deep-marine (upper-middle bathyal) claystone.

952 **Fig. 2.** Seabed topography, as defined by this bathymetry map, shows the external geometry of the
953 Haya Slide. The slide originated from the NE (collapse of the southern flank of a thrust-cored anticline)
954 and transported towards the SW. This study focuses on the toe domain of the slide (red outline), which
955 is mostly imaged by the 3D seismic reflection data (blue outline). The toe domain of the slide has a
956 radial geometry, where the Eastern and Northern lateral margins trending N-S and E-W, respectively.

957 **Fig. 3. (a)** Thickness map covering the toe domain of the Haya Slide. The slide is thickest (200 m) in the
958 southern part and thins toward the Northern Lateral Margin. Laterally, three areas can be defined
959 based on its frontal geometry (i.e. Area A, B, and C). An inset map showing the focus area of the slide,
960 captured by 3D seismic reflection data. **(b)** Spectral decomposition map showing internal seismic
961 facies of the slide. Axially, the slide can be divided into inner and outer parts with 'soft' boundary
962 between them. The inner part is dominated by debrite containing megaclasts, and the outer part is
963 dominated by pop-up blocks.

964 **Fig. 4.** Seismic sections across Area A, B, and C, showing similar general characteristics, where debrite
965 dominates the inner part, and pop-up blocks dominate the outer part. However, the three areas have

different characteristics of frontal margin. **(a)** Area A is characterised by frontal confinement and coherent pop-up blocks. Translation distance was estimated by calculating shortening amount at H1 and 2, i.e. 8-14% shortening equating to 0.6-1.1 km. **(b)** Area B is characterised by frontal ramps with more chaotic reflections adjacent to frontal margin, and less coherent pop-up blocks. **(c)** Area C is characterised by frontal emergence and a broad bulge on the seabed above steeply-inclined detachment surface.

Fig. 5. Deformation ahead of the parent flow. **(a)** Variance time-slice showing distributed shear zone downdip from an intact block. Thrusts forming this distributed shear zone laterally propagate eastwards. **(b)** Seismic section showing distributed shear zone, showing deformed strata ahead immediately downdip from the intact block. Folded strata ahead of the BSS, interpreted as an unformed thrust.

Fig. 6. (a) Basal shear surface structure map showing slope gradient break in Area C. **(b)** Seabed structure map showing a broad area of high seabed relief (seabed bulge). **(c)** Spatial relationship between slope gradient break on the BSS and the occurrence of the seabed bulge, leading to frontal emergence of the slide. **(d)** Seismic section adjacent to Northern Lateral Margin showing closely-spaced pop-up blocks and frontal confinement of the slide.

Fig. 7. (a) Variance along the BSS (50 ms windowed above) showing an abrupt boundary between Area B and C. **(b)** A ramp marks the boundary between Area B and C, and expressed as positive relief on the seabed.

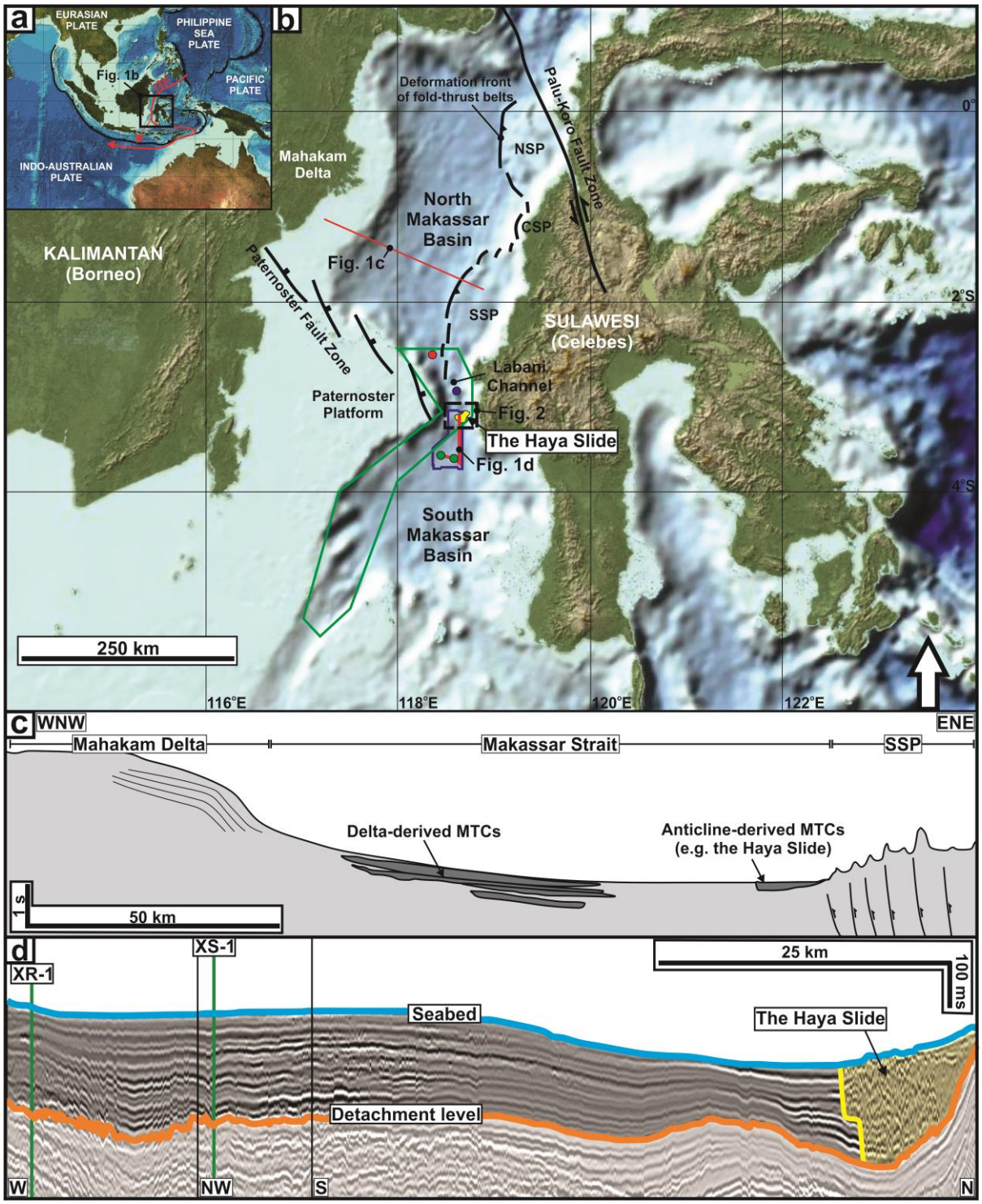
Fig. 8. Along-strike quantitative analysis of Pop-up Block 3 (see Fig. 4a). **(a)** Time structure map of H2 (see Fig. 4a) and associated faults. **(b)** Variance time-slice showing lateral extent of Pop-up Block 3. **(c)** Throw vs. Distance (T-x) plot of fore- and backthrusts bounding Pop-up Block 3. Shear zone separates two bodies that have different amount of strain, i.e. the area to the west of the shear zone experienced more contraction as shown by cumulative throw as compared the area eastwards from the shear zone. **(d-f)** Seismic sections showing along-strike variability of faults bounding Pop-up Block 3.

Fig. 9. Schematic model of emplacement processes of the Haya Slide. **(a)** Debris flow, originated from failed anticline (see Fig. 2) entered the basin, deformed the seabed, and then entrained substrate into the flow. **(b)** Substrate erosion and entrainment continued to occur up to the point where the debris flow did not have sufficient shear stress for substrate entrainment. Thus, the remaining exerted stress deformed substrate ahead of the flow (i.e. lateral loading). **(c)** Subsequent compressional deformation occurred, allowing a relatively short translation distance (0.61 to 1.06 km) in the toe domain, which has different frontal geometries along strike.

Fig. 10. A summary of downdip and along-strike variations in Areas A, B and C of the Haya Slide., Note the lateral changes in structural style and internal facies characteristics.

Fig. 11. Evolution of basal shear surface adjacent to the toewall of the Haya Slide, showing development of **(a)** frontal confinement in Area A, **(b)** frontal emergence in Area C, and **(c)** staircase-like frontal ramps in Area B, which is an intermediate (transitional) style between frontal confinement and emergence.

1005 Figure 1



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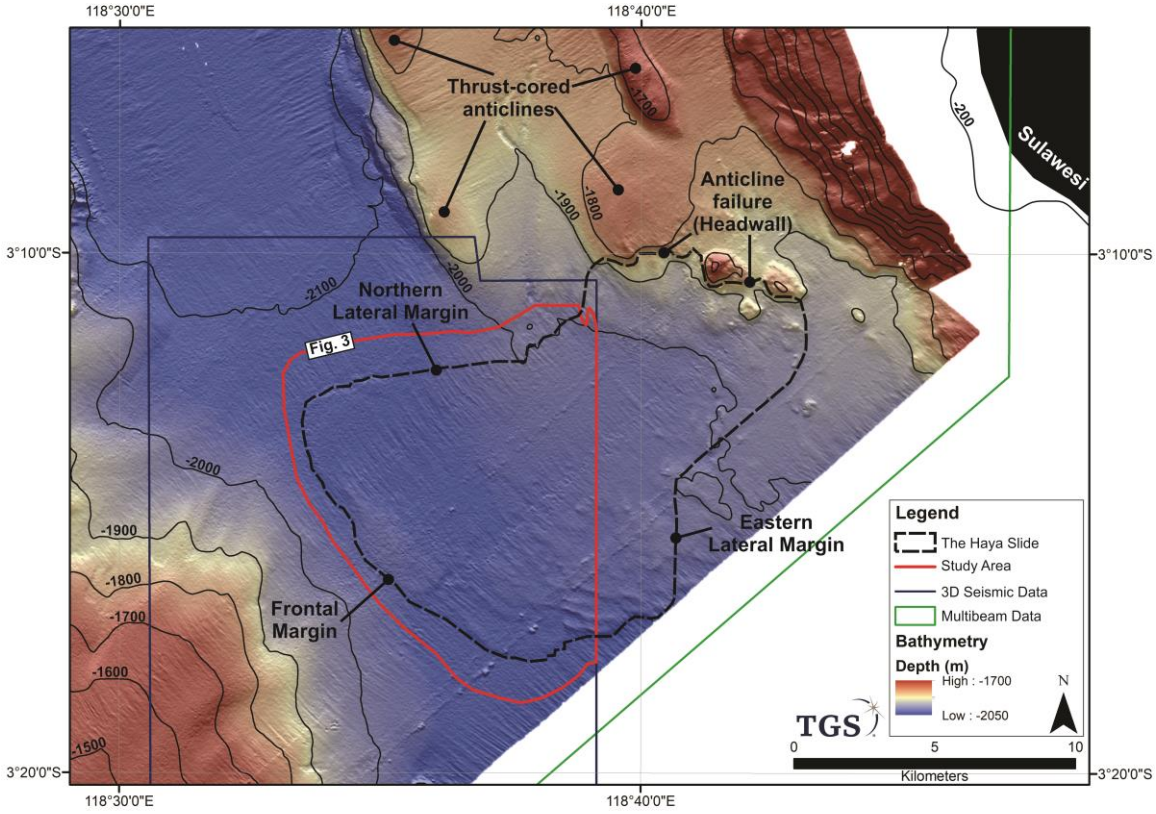
1007 Figure 2



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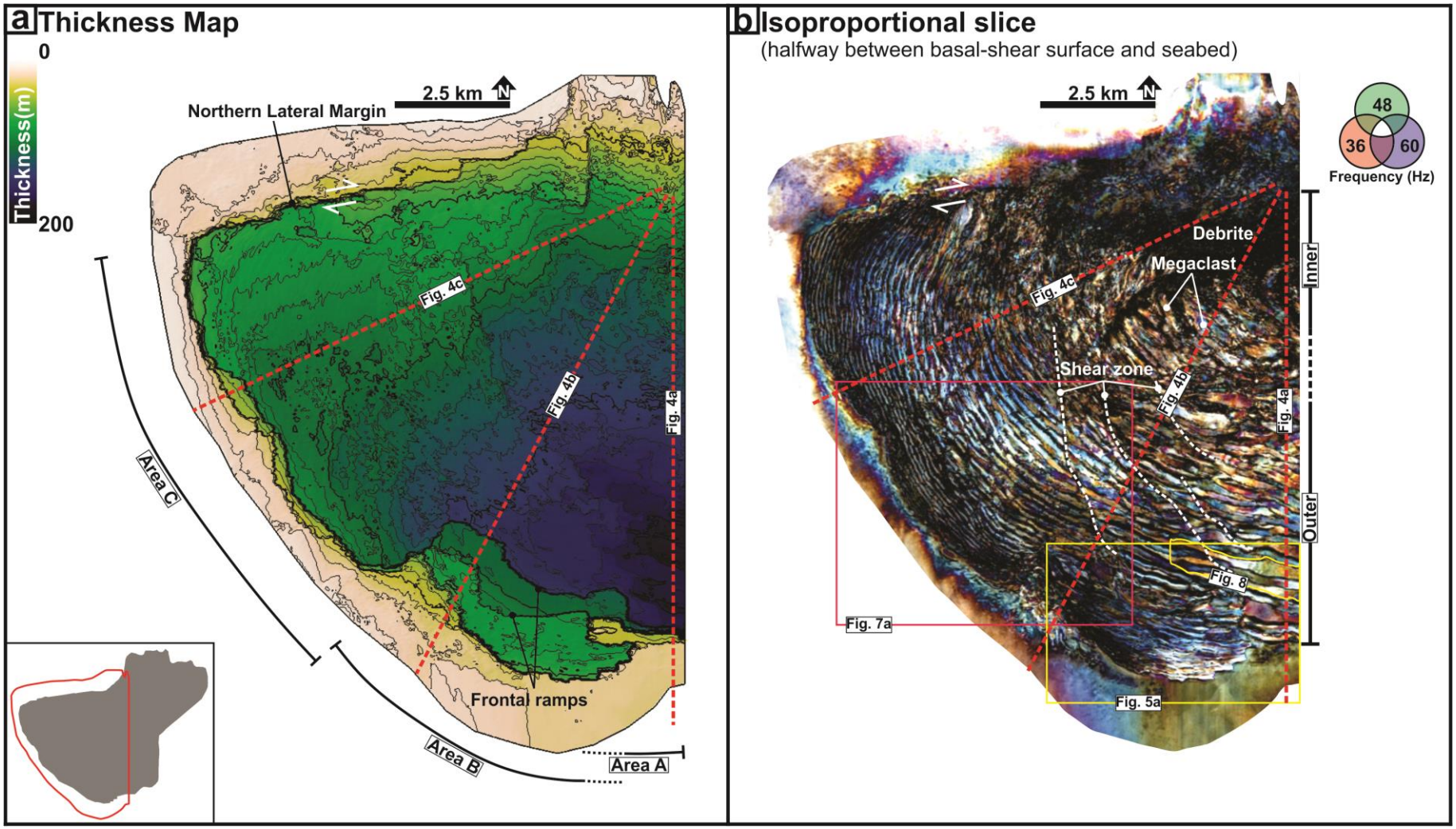

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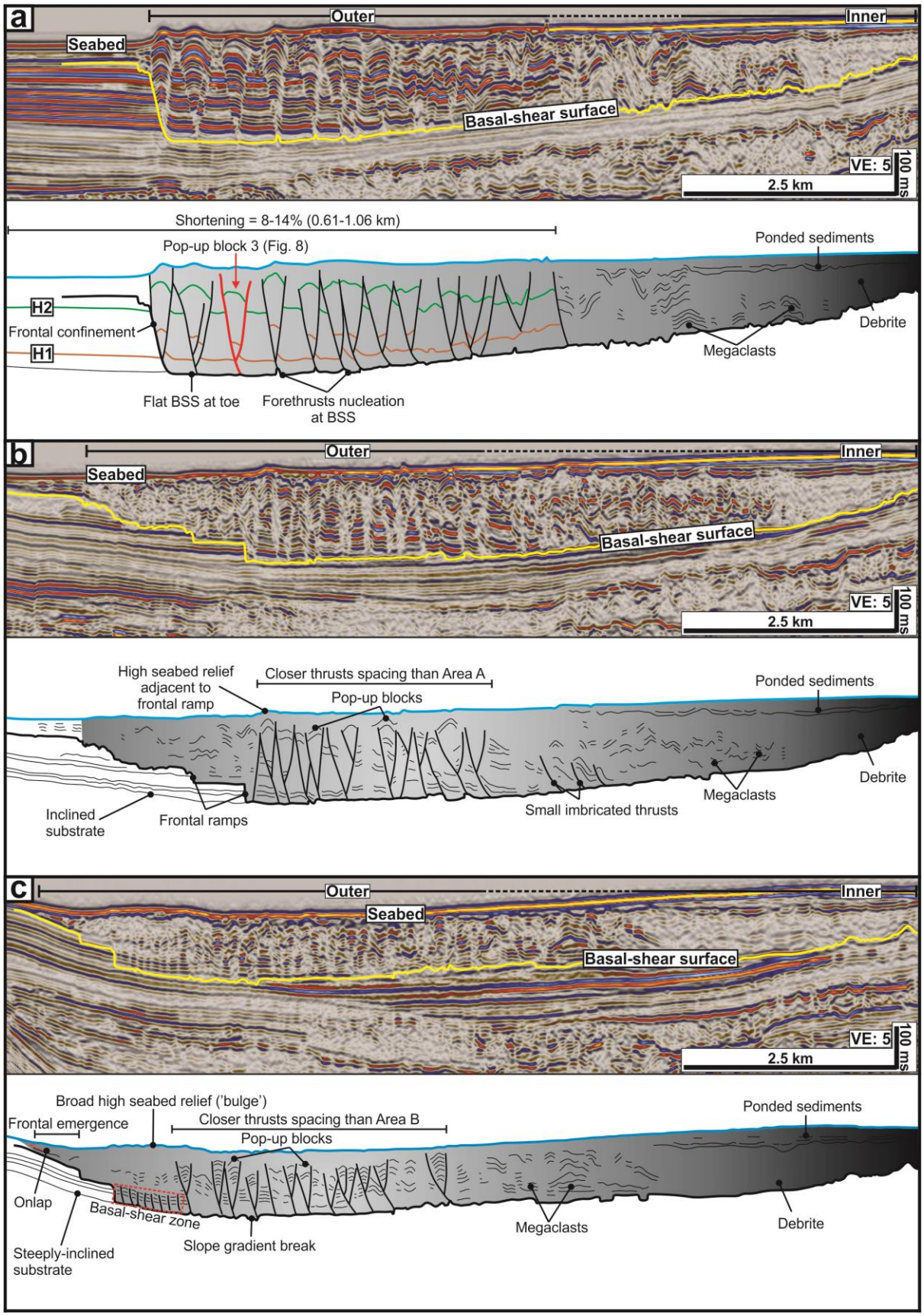

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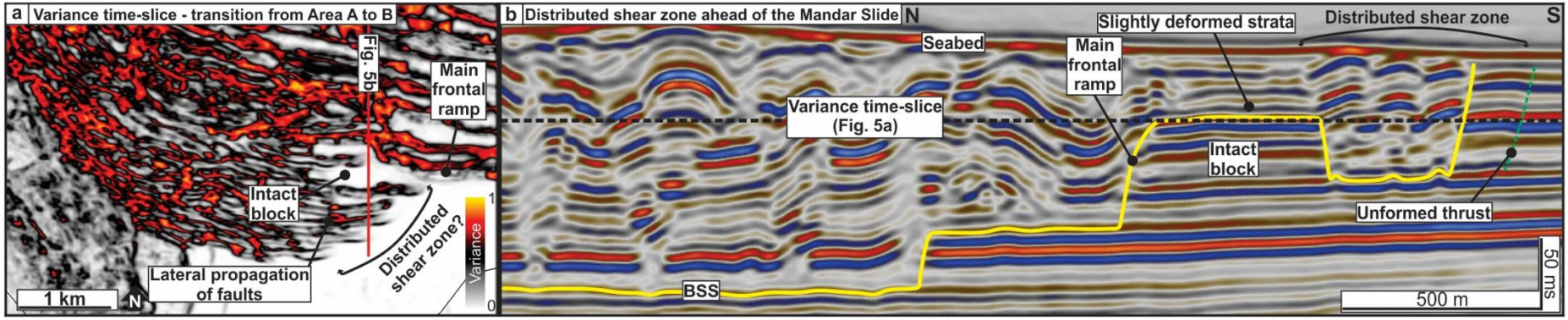

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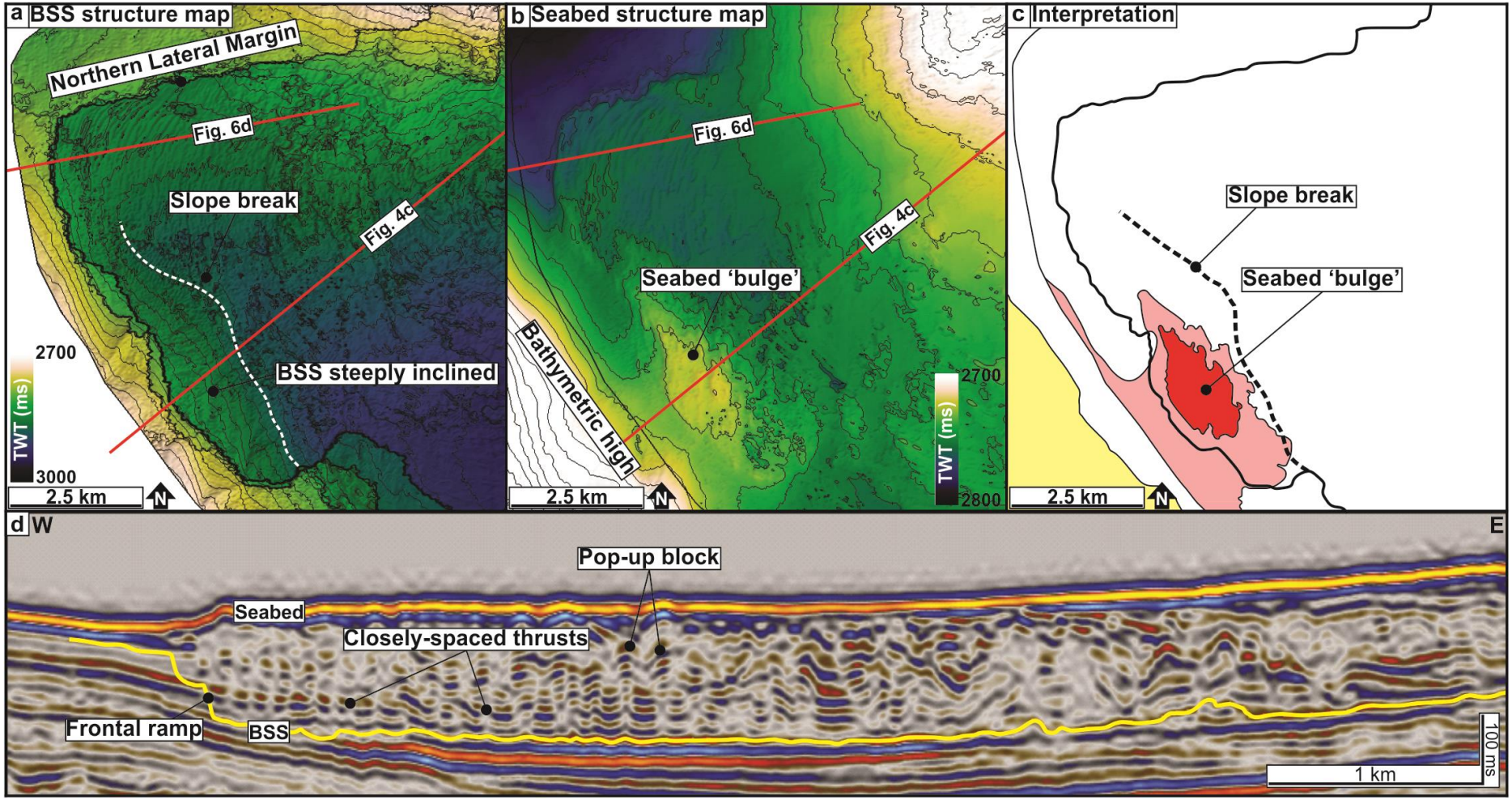

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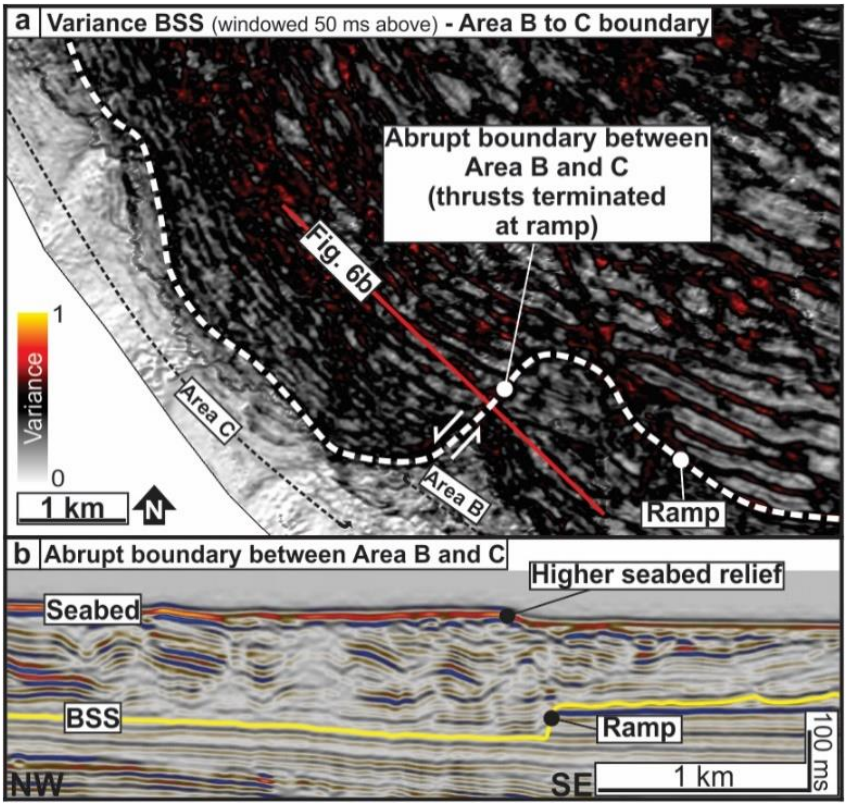

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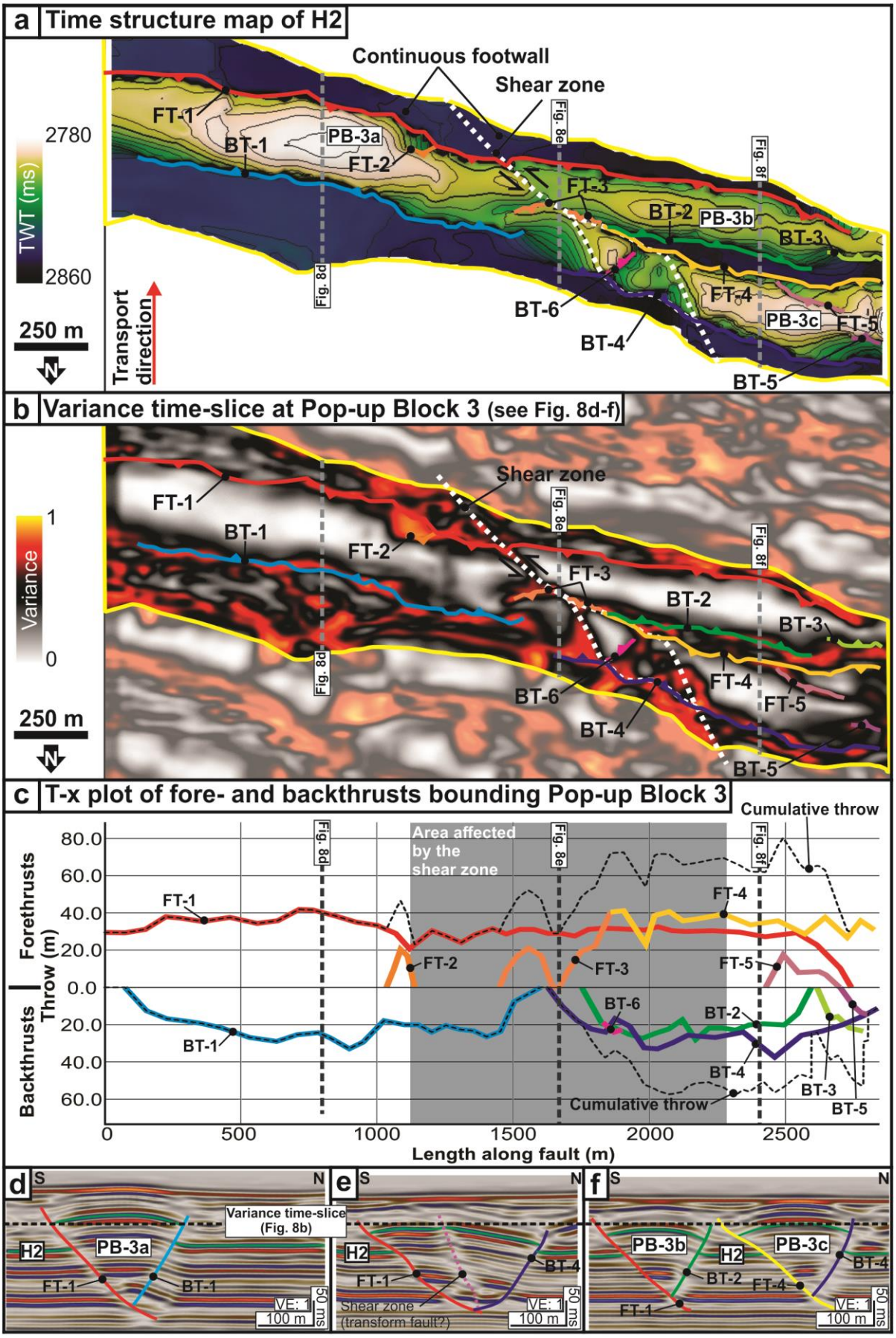

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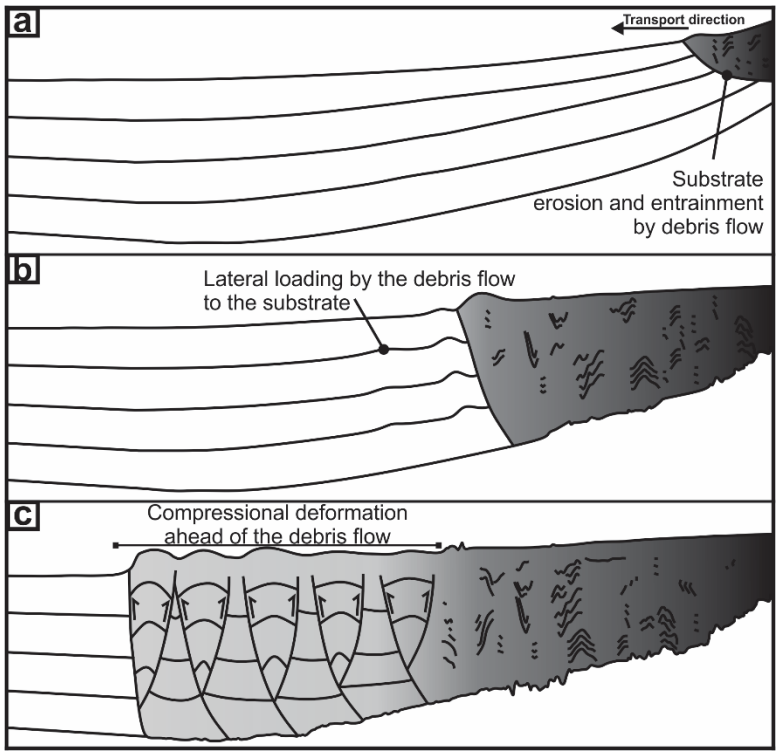
1021 Figure 9



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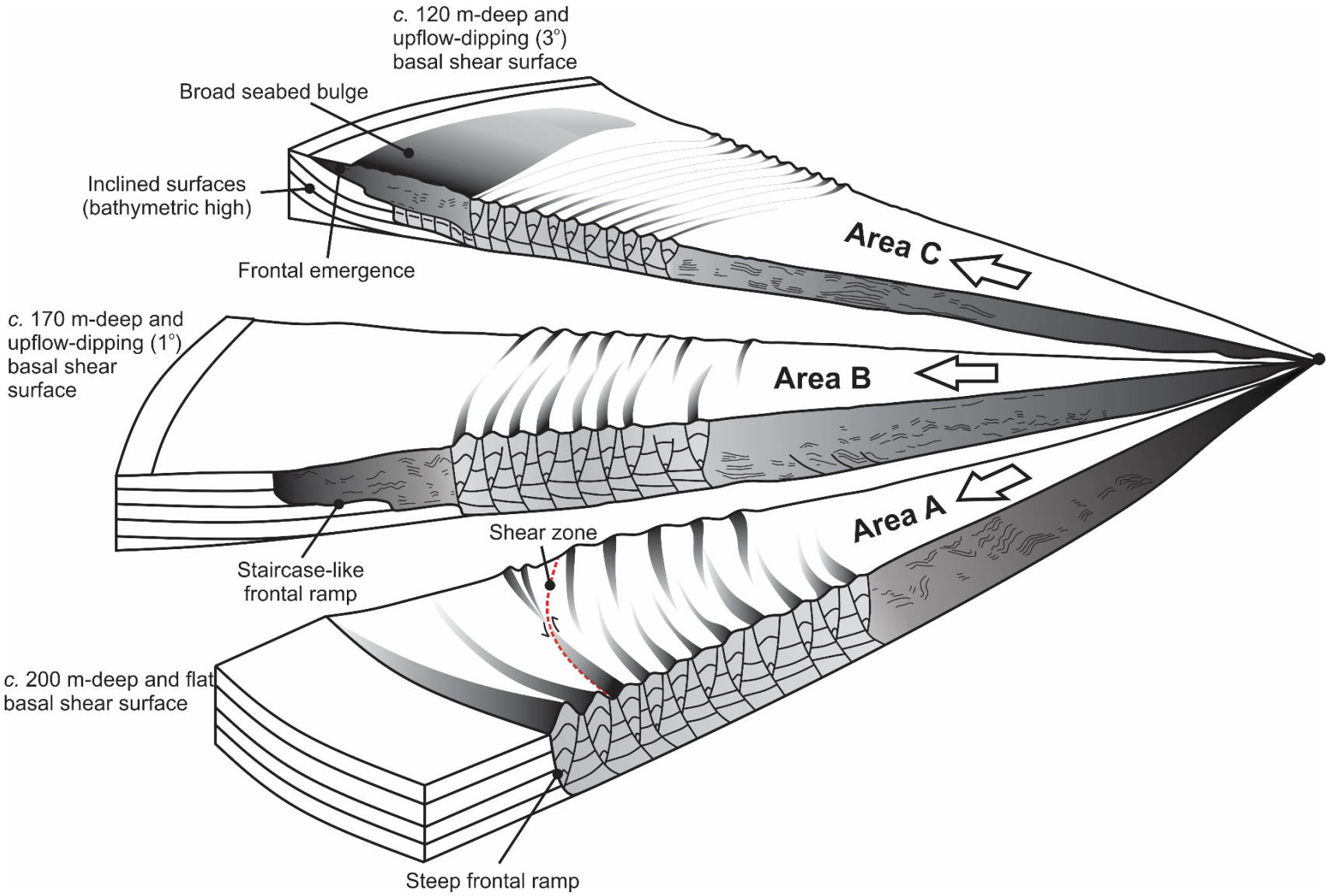

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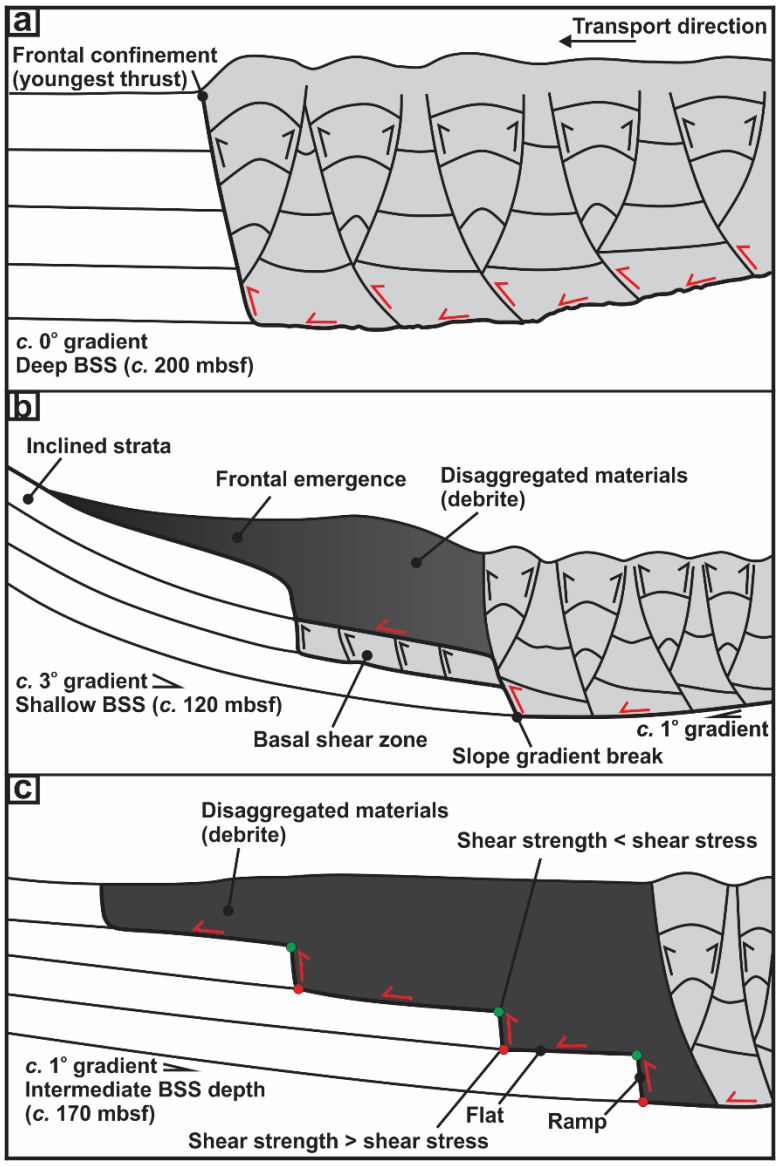
1025 Figure 11



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