**The effects of a weak asthenospheric layer on surface kinematics, subduction dynamics and slab morphology in the lower mantle**

# The effect of a weak asthenospheric layer on surface kinematics, subduction dynamics and slab morphology in the lower mantle

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## Abstract

On Earth, the velocity at which subducting plates are consumed at their trenches (termed 'subduction rate' herein) is typically 3 times higher than trench migration velocities. The subduction rate is also 5 times higher than estimated lower mantle slab sinking rates. Using simple kinematic analyses, we show that if this present-day "kinematic state" operated into the past, the subducting lithosphere should have accumulated and folded beneath near-stationary trenches. These predictions are consistent with seismic tomography, which images localized and widened lower-mantle slab piles. They are, however, at odds with most dynamic-subduction models, which predict rapid trench retreat and inclined slabs in the mantle transition zone. We test the hypothesis that a weak asthenospheric layer (WAL), between the lithosphere-asthenosphere boundary and 220 km depth, compatible with geophysical constraints, can remedy the discrepancies between numerical models and observations. The WAL lubricates the base of the lithosphere, increases the subduction rate while reducing trench retreat. As a consequence, simulations featuring a WAL predict slab accumulation at the mantle transition zone, and thicker, folded slabs in the lower mantle. A WAL viscosity only 2-5 times lower than that of the adjacent mantle is sufficient to shift subduction regimes towards a mode of vertical slab sinking and folding beneath near-stationary trenches, across a wide range of model parameters, producing surface and slab velocities close to those observed at the present-day. These findings provide support for the existence of a weak asthenosphere beneath Earth's lithosphere, complementing independent evidence from various geophysical data.

# 1 Introduction

The negative buoyancy of subducting plates is the primary driving force sustaining subduction and surface plate motions Forsyth & Uyeda (1975). Subduction zones are the sites of tectonically-forced horizontal deformation Uyeda & Kanamori (1979); Lallemand et al. (2005) and dynamic vertical motions Davies (1981); Gurnis (1993). Crust and lithosphere subducting beyond the mantle transition zone add chemical heterogeneities to the lower mantle, which are stirred and homogenised by mantle convection Zindler & Hart (1986); Jones et al. (2016), or persist to the core-mantle-boundary, as suggested by modern tomographic models (e.g. Hosseini et al., 2020). Understanding the deep dynamics of subducting slabs is thus key for addressing the geodynamical and geochemical evolution of our planet.

Observed plate kinematics provide insights into the dynamics of the subduction system Forsyth & Uyeda (1975); Jarrard (1986); Lallemand et al. (2005); Heuret & Lallemand (2005); Sdrolias & Müller (2006); Doglioni et al. (2007); Funiciello et al. (2008); Schellart (2008b); Becker & Faccenna (2009); Goes et al. (2011). Subduction kinematics (see Fig. 1) involve the velocities of the subducting plate $v_{sp}$ ("SP velocity" for short); the velocity of the overriding plate $v_{op}$ ("OP velocity"); and the velocity of the trench $v_t$, which is equal to OP velocity if the overriding plate does not undergo (back-arc) deformation. Note that $v_{sp}$ and $v_t$ are defined with opposite signs: the natural (positive) direction of trench migration is "retreat" towards the SP. These velocities are given in some absolute reference frame, which is taken as the stable lower mantle herein Becker & Faccenna (2009).

We also use a relative velocity, the subduction rate $v_s$. This is the velocity of the subducting plate relative to the trench (i. e., the rate at which the subducting plate is consumed by the migrating trench). It has been repeatedly shown that typical values of $v_s$ on Earth are higher than 3-4 cm/yr Forsyth & Uyeda (1975); Jarrard (1986), while absolute trench motions are usually between -2 and 2 cm/yr Heuret & Lallemand (2005); Funiciello et al. (2008); Schellart (2008b). Other studies have pointed out that the magnitude of the (absolute) SP velocity $v_{sp}$ is generally two to three times higher than that of the (absolute) $v_t$ Becker & Faccenna (2009); Goes et al. (2011); Carluccio et al. (2019). Hence plates are consumed at much faster rates than their trenches move laterally.

Analogue and numerical models of subduction dynamics without external forcing (hereafter simply referred to as models of subduction dynamics) have shed light on the internal force balance of subduction systems and the resulting kinematics. They have illuminated various subduction regimes and slab morphologies in the upper mantle (e. g. Guillou-Frottier et al., 1995; Schellart, 2008a; Di Giuseppe et al., 2008; Ribe, 2010; Stegman et al., 2010). Recent studies that included an overriding plate with finite strength, concluded that the slab pull force associated with the negative buoyancy of a subducting plate (SP) favored slab rollback and migration of of the trench towards the subducting plate (i. e., trench retreat), unless the SP was weak and/or the overriding plate (OP) was strong Garel et al. (2014); Sharples et al. (2014); Holt et al. (2015); Hertgen et al. (2020). It has been pointed out that such analogue and numerical models of subduction dynamics tend to produce surface kinematics that are at odds with some of the first-order observations outlined above Goes et al. (2011); Carluccio et al. (2019). These subduction models generally produce trench retreat velocities that exceed present-day observations, especially once the subducting slab reaches the bottom of the upper mantle, which was sometimes treated as a rigid barrier Funiciello et al. (2004); Schellart (2005);

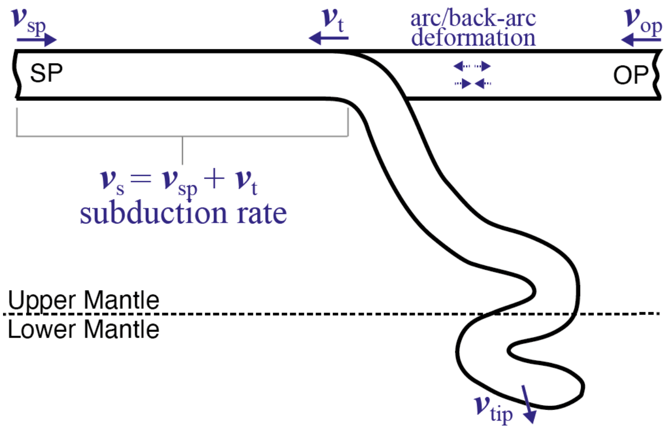


Figure 1: a) Sketch illustrating the various kinematics that can be inferred within a subduction system. Positive values denote absolute trenchward motion for both the subducting plate ($p$) and the overriding plate ($v_{op}$). Absolute trench velocity ($v_t$) is considered positive towards the subducting plate. We define relative to a fixed lower-mantle reference frame, in order to compare to absolute motions in nature, which can be quantified in empirical, approximate mantle reference frames, such as the fixed-hotspot frame. Subduction velocity ($v_s$) is a relative velocity, the rate at which the subducting plate is consumed at the trench ($v_s = v_{sp} + v_t$).

Capitanio et al. (2007); Goes et al. (2011). More modest trench motions over a relatively wide range of parameters have only been produced by 2-D models that consider both the penetration of the subducting slab into the lower mantle and complex rheologies Garel et al. (2014); Holt et al. (2015); Li et al. (2019). Even then, slower trench motion is only achieved at the cost of decreasing the SP velocity to values of less than 2 cm/yr once the slab interacts with the viscosity increase around 660-km depth (hereafter referred to as "first slab-660 interaction") (e.g. Garel et al., 2014; Suchoy et al., 2021). Hence, in most subduction dynamics models, more than half of the subduction rate $v_{sp}$ is accounted for by trench motion $v_t$, which contradicts present-day observations of plate kinematics.

Independent constraints on subduction dynamics come from seismic tomographic images of slab morphologies at and below the mantle transition zone. A few slabs under present-day subduction zones in the Western Pacific appear to stagnate above the 660-km discontinuity Karason & Van Der Hilst (2000); Amaru (2007); Li et al. (2008); Fukao & Obayashi (2013) - for instance, under Japan Fukao et al. (1992) and under Izu-Bonin Wu et al. (2016), at least under its northernmost part Zhang et al. (2019). But many other slabs have breached the 660-km discontinuity and are sinking into the lower mantle Goes et al. (2017). Transition-zone and lower-mantle slabs are imaged more robustly and consistently than slabs in the upper(most) mantle. The opposite would be expected if slabs retained a constant thickness across depths. Hence the deeper slab must be thicker Ribe et al. (2007); Loiselet et al. (2010), which is well-documented under the Americas Karason & Van Der Hilst (2000); Ren et al. (2007); Sigloch & Mihalynuk (2013); Mohammadzaheri et al. (2021), but also globally Van der Voo et al. (1999); Shephard et al. (2017); Van der Meer et al. (2018); Hosseini et al. (2020). Under the particularly well-instrumented Cascadia subduction zone of North America, tomography can resolve a shallow slab of single lithospheric thickness, and also confidently show that the slab is multiply thickened from the transition zone downward Sigloch et al. (2008).

Thickened slabs in the lower mantle have been attributed to slab buckling and folding through the mantle transition zone Guillou-Frottier et al. (1995); Ribe et al. (2007); Běhounková & Čížková (2008); Lee & King (2011); Cerpa et al. (2014); Billen & Arredondo (2018), with possible slab detachment Čížková et al. (2012). Slab folds have not yet been resolved by tomography, so the exact widening mechanism remains speculative from the observational side.

In order to produce slab folding and realistic lower-mantle slab morphologies, numerical subduction models often require a fixed overriding plate, i.e. zero trench velocity $v_t$. Even then, the models predict sub-vertical slab folding only for rather extreme values of model parameters, e. g. a very young subducting plate Garel et al. (2014); Strak & Schellart (2021), and/or if special model setups are considered such as those involving two nearby subduction zones Čížková & Bina (2015); Lyu et al. (2019). Thus the observations of pervasively thickened lower-mantle slabs are generally not predicted in current models of subduction dynamics. The models may lack a first-order mechanism that generates thick lower-mantle slabs.

To summarize, there are at least two discrepancies between existing models of subduction dynamics and first-order observations. First, current models generally produce trench retreat velocities $v_t$ in excess of those observed at present-day subduction zones, alongside SP velocities $v_{sp}$ and subduction rates $v_s$ that are too slow after first slab-660 interaction. Second, models seldom reproduce the tomographically observed, multiply thickened geometries that prevail in the transition zone and lower mantle.

This study considers how a weak asthenospheric layer (WAL) beneath the plate can resolve these discrepancies. The presence of a WAL on Earth has been proposed to explain a large range of geophysical observations, including postglacial rebound and gravity data (e.g. Paulson & Richards, 2009), shear-wave tomography Kawakatsu et al. (2009); Barruol et al. (2019), seismic attenuation Debayle et al. (2020), seismic anisotropy Debayle & Ricard (2013); Becker (2017) and electrical conductivity tomography Naif et al. (2013). The viscosity reduction could originate from a plume-fed asthenosphere Phipps Morgan et al. (1995), from the depth-dependency of dislocation creep flow laws Raterron et al. (2011), from crystal-preferred orientation Meyers & Kohlstedt (2021), or from the presence of melt pockets Cooper & Kohlstedt (1986); Chantel et al. (2016), which may remain trapped due to low melt fractions Holtzman (2016) or low density contrast Sakamaki et al. (2013).

The presence of a WAL is predicted to affect large-scale dynamics of the underlying, convecting mantle Lenardic et al. (2006), and to favor 'plate-like' rather than 'stagnant-lid' regimes Höink et al. (2012). Since the sub-lithospheric mantle resists a plate's trenchward motion, the inclusion of a WAL in models of subduction dynamics yields faster subduction velocities $v_{sp}$, as shown by Carluccio et al. (2019) and Suchoy et al. (2021). The latter authors also showed that increased $v_{sp}$ was coeval with reduced trench retreat $v_t$, although they did not detail the implications for lower mantle slab morphologies. We hypothesize that increasing subduction rates while reducing trench motion results in the accumulation of slab material in a near-vertical column beneath the (quasi stationary) trench, and that the slab must widen (through folding) around the depths where it slows down to lower-mantle sinking rates, given that slab input $v_s$ remains high. Thus, a WAL could resolve both first-order discrepancies regarding plate velocities and slab morphologies.

We carry out a systematic numerical analyses of how a WAL impacts the dynamics of thermo-mechanical subduction models featuring an overriding plate. Section 2 provides a first-order quantification of slab widening behavior in modern subduction zones, using plate kinematic data. Section 3 describes the model setup. Section 4 presents our modeling results, and Section 5 discusses their implications for subduction systems on Earth.

# 2 Quantifying slab folding from plate motions and slab sinking rates

## 2.1 Conceptual assessment

We start by demonstrating how slab folding can be assessed theoretically as a geometrical/kinematic phenomenon, involving slab accumulation in the mantle transition zone. This analysis is inspired by subduction models where (unlike our own model) velocities are applied to one or both plates, and which can predict slab morphology as a function of these imposed surface kinematics Christensen (1996); Heuret et al. (2007); Arcay et al. (2008); Gibert et al. (2012); Cerpa et al. (2015); Guillaume et al. (2018); Cerpa et al. (2018). Among these, Gibert et al. (2012), anchored the subducting slab to a rigid 660-km discontinuity, which aims to simulate the effect of a strong viscosity jump at 660 km, as inferred from e.g. geodetic constraints Mitrovica & Forte (2004). Gibert et al. (2012) showed that if the subduction rate $v_s = v_{sp} + v_t$ exceeds the trench velocity $v_t$, continued subduction results in slab folding at the base of the upper mantle. Essentially, the slab has to fold because trench retreat does not create enough lateral accommodation space to permit all incoming slab to lie down flat on the '660'. Here we extend their analysis to the more general case where the subducting slab sinks into the lower mantle.

Let $v_s \times \Delta t$ be the length of subducted material consumed at the trench over some duration $\Delta t$. The lateral displacement of the trench over the same duration is $v_t \times \Delta t$. The displacement of the deepest portions of the subducting slab (simplified as the displacement of the slap tip) within the upper mantle is approximated as $v_{\text{tip}} \times \Delta t$, where $v_{\text{tip}}$ is the absolute velocity of the deepest point of the slab. Slab folding can thus be understood as a simple geometrical constraint. When the length of subducted material is larger than the lateral displacement of the trench plus the displacement of the slab tip, the excess length (slab accumulation) is expected to be accommodated by folding. Put in another form, slab accumulation and folding occurs when:

consumption rate of subducted material > trench displacement + slab tip motion in the mantle

Alternatively, we can define a kinematic ratio $K_r$ which predicts whether the subducting slab undergoes folding as:

$$K_r = \frac{v_s}{|v_t| + \sqrt{(v^x_{\text{tip}})^2 + (v^z_{\text{tip}})^2}} \tag{1}$$

where we have decomposed the velocity of the deepest point of the slab into its horizontal and vertical components.

When $K_r \simeq 1$, the free space created by trench retreat and slab sinking can accommodate all newly incoming lithosphere, which does not have to compress (fold). Hence the slab's apparent thickness remains similar in the upper and lower mantle (Fig. 2). A kinematic ratio $K_r$ higher than 1 implies a surplus of slab material that cannot be accommodated by trench retreat and slab sinking into the lower mantle, and instead has to be accommodated by slab folding. At $K_r > 1$, the higher the value of $K_r$, the greater the frequency of slab folds (or alternatively, the wider the amplitude of the folds). Also at $K_r > 1$, the apparent thickness of the folded slab in the lower mantle is predicted to be multiples of the lithospheric thickness observed in the uppermost mantle.

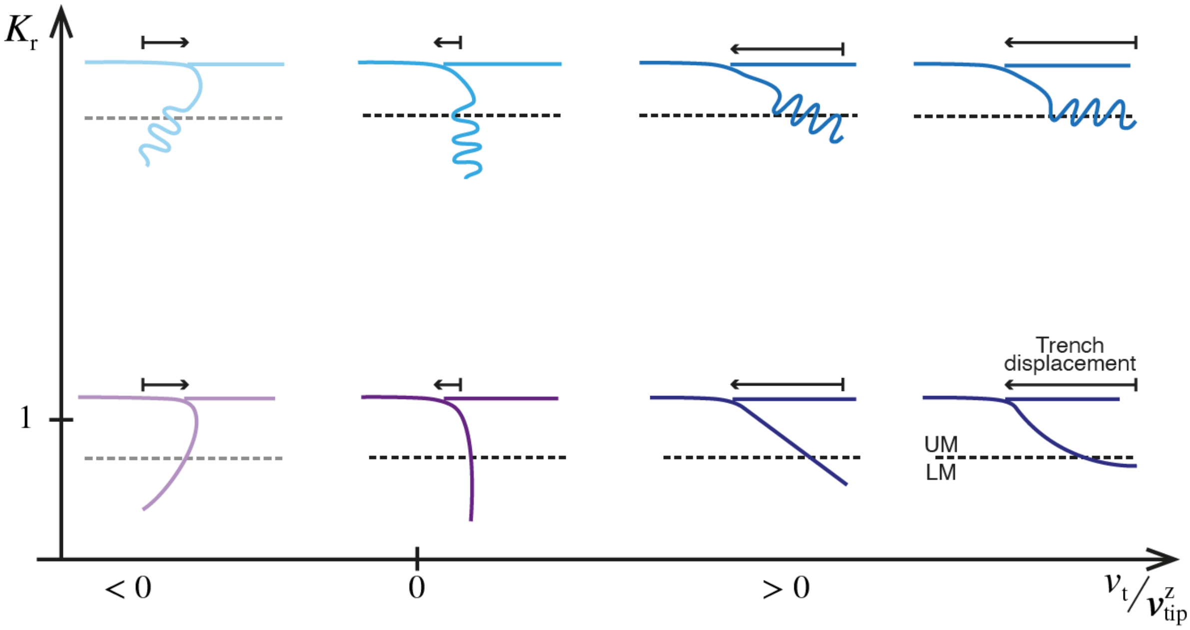


Figure 2: Schematic theoretical diagram depicting various subduction regimes of subducting slabs, well after they have penetrated in to the lower mantle (LM). UM: upper mantle. Dashed line marks the viscosity interface between UM and LM, which we speculatively equate with the seismic '660 km' discontinuity.

The ratio of trench velocity to slab-sinking velocity ($v_t/v_{\mathrm{tip}}^z$) controls, to first order, the average slab dip. This is true for both the unfolded ($K_r \leq 1$) and folded ($K_r > 1$) cases. Note that we use only the vertical component of the slab-tip velocity as it is thought to be much higher than the horizontal component (see also below section 2.2).

Hence the parameter space along the dimensions of $K_r$ and $v_t/v_{\mathrm{tip}}^z$ spans a variety of candidate slab morphologies and subduction regimes, as depicted in (Fig. 2). For negative trench motion (trench advance), the subducting slab leans forward, so that the deepest slab portions lie beneath the subducting plate at increasing distances from the trench (leftmost regimes in Fig. 2). For quasi-null trench motions, the subducting slab sinks vertically, with all slab portions remaining below the trench. For relatively high, positive trench motions (trench "retreat"), the subducting slab leans backwards (" slab rollback"), with deeper slab beneath the overriding plate. It has been proposed that high trench-retreat rates promote the complete stagnation of slab atop the 660-km discontinuity Torii & Yoshioka (2007); Goes et al. (2017), so that high values of $v_t/v_{\mathrm{tip}}^z$ may lead to the end-member subduction regime where the slab flattens and folds on the 660-km discontinuity (rightmost regimes in Fig. 2).

## 2.2 Estimating slab folding at present-day

In order to gauge the prevalence of slab folding in nature, we seek to calculate an observational estimate of $K_r$ in active subduction zones, using Equation 1. Hence we need estimates of subduction rate $v_s$, absolute trench velocity $v_t$, and slab sinking velocity $v_{\mathrm{tip}}$.

For estimating $v_s$ and $v_t$, we use an updated version of the SUBMAP database Lallemand et al. (2005), which defines 249 transects of active subduction zones. Subduction rates are retrieved from the relative plate motions of the MORVEL56-NNR model (based on a circuit of 56 tectonic plates Argus et al. (2011) as explained in the Supporting Information Text S1). Each SUBMAP subduction transect is assigned to a subducting plate and an overriding plate of the MORVEL56-NNR plate circuit. For transects that cross significant arc and back-arc deformation, MORVEL56-NNR permits the definition of an "arc block" and assessment of trench motion relative to that of a rigid overriding plate, enhancing the accuracy of the derived subduction rate. For a few subduction zones, the MORVEL56-plate circuit does not account for active arc and back-arc deformation even though such a deformation has been well-established in the literature (Southernmost-Central Andes, Izu-Bonin, Calabria). For these transects, we complement MORVEL56-NNR with published regional studies (see Supporting Information).

To define the absolute motion of the plates and trenches, we need to consider an absolute plate motion model within an absolute reference frame, comparable to the fixed reference frame of our numerical models. In this paper we calculate and compare the value of $K_r$ in three recent absolute plate motion models, constructed in different manners: the "SA" ("spreading-alignement") model Becker et al. (2015), the "TM25" model Wang et al. (2018), and the "GMHRF-1Ma" model by Doubrovine et al. (2012). The SA model minimizes the angular misfit between spreading-ridge orientations and plate velocities. This plate motion was found to give a good fit to azimuthal seismic anisotropy, a proxy for the shear induced by the relative motion between the tectonic plates and the upper mantle. The TM25 model is based on 25 hotspot tracks under the assumption of fixed hotspots relative to the deep mantle. The GMHRF-1Ma model is based on a global fit of hotspot tracks since the Late Cretaceous, accounting for modest relative motions between the hotspots' mantle plumes, computed by numerical models of whole mantle convection.

We extract the trench-normal component of the plates and trench velocities for comparison with our 2-D models. In what follows, the absolute and relative velocities at each transect are those of their trench-normal components.

The observed subduction rates are non-negative with a median value of 5.3 cm/yr and a long tail up to almost 12 cm/yr (Fig. 3a). In all three reference frames (Fig. 3b-d), absolute trench velocity $v_t$ scatters around slightly positive values with a median values of 0.71 to 0.79 cm/yr. This tendency towards slow trench retreat may or may not be significantly different from zero motion (stationary trench), given the large formal standard deviations of almost 3 cm/yr but also the non-Gaussian, heavy tails of the histogram. In any case, two thirds of the subduction transects have trench velocities between -1 and 1 cm/yr in the three absolute plate motion frames. Hence typical present-day trench motion is roughly five times smaller than typical subduction velocities.

Estimating $K_r$ also requires an observational estimate of slab sinking rates in the lower mantle. Since slabs are not directly dateable, they have been correlated to the geology of accretionary orogens, which hold the surface record of subduction. The subduction of lithosphere is accompanied by the formation of a volcanic arc at the surface, which often survives and is dateable. Such slab-arc correlations have inferred time-averaged sinking rates of 1.0-1.5 cm/yr for slabs that have penetrated the lower mantle Van Der Meer et al. (2010); Sigloch & Mihalynuk (2013); Domeier et al. (2016); Van der Meer et al. (2018); Mohammadzaheri et al. (2021).

Using 1 cm/yr as the slab sinking velocity estimate, 70–80% of subduction transects exhibit values of $K_r > 1$ in all three absolute reference frames (Fig. 4 and Fig. S1). Only

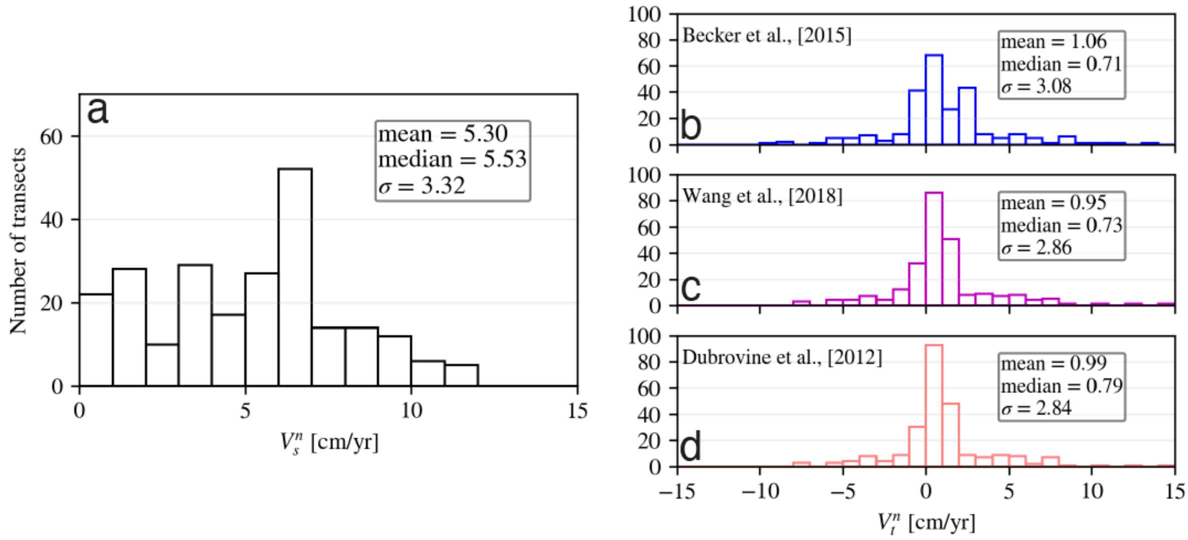


Figure 3: a) Histogram of the trench-normal component of subduction rate $v_s$ in present-day subduction zones. b Histogram of the trench-normal component of trench velocity (retreat) $v_t$, in the spreading-alignment reference frame Becker et al. (2015). c) Histogram as in b) but for T25M reference frame Wang et al. (2018). d) Like b), but for GMHRF-1Ma frame Doubrovine et al. (2012).

a few subduction transects consistently display $K_r < 1$ in all reference frames, mostly at the edges of longer arcs: the southernmost Andes (Patagonian transects), the northern edge of the Lesser Antilles (e.g. Puerto Rico Trench), or the edges of the South Sandwich SZ.

The present-day prevalence of $K_r > 1$ is relatively insensitive to the assumed slab sinking velocity. Even when considering $v_{tip} = 1.5$ cm/yr, at the high end of the reasonable estimate range ?see e.g.>[]butterworth2014geological,domeier2016global), 63% of transects remain above $K_r > 1$ in the spreading-aligned absolute plate motion model, and 72% of transects in the two other reference frames.

Figure 4a plots the global inventory of slabs (between 600-1800 km depth), from which the sinking rates were derived. Importantly, most areas are slab free. Existing slabs cluster in two vast, linear belts: one under the Alpine-Eurasian-Himalayan-southwest Pacific orogens; the second under the Americas and into Siberia. From the geologic record and quantitative plate reconstructions, these are the known, absolute locations of major orogenies over the past 200 million years, hence the known paleo-trench locations. The observation that slabs are still located only beneath these independently inferred paleo-trench regions means that slabs sank rather vertically. The vast slab-free mantle areas are known not to have hosted trenches over the past 200 m.y. This implies that paleo-trenches have remained quite stationary over a time period during which the areal equivalent of all ocean basins was subducted once or twice over Coltice et al. (2012). Thus trenches had the opportunity to migrate across the globe but did not, which indicates sustained $K > 1$ (slab folding regime) over geologic time.

Finally, slab dimensions directly point towards folding. In figure 4a, the Eurasian and American slab belts are 15,000-20,000 km long; individual slab segments are 1,000-3,000 km long (i.e., arc length) and 400-700 km wide. The latter is a multiple of lithospheric thickness, and suggests slab folds of this amplitude.

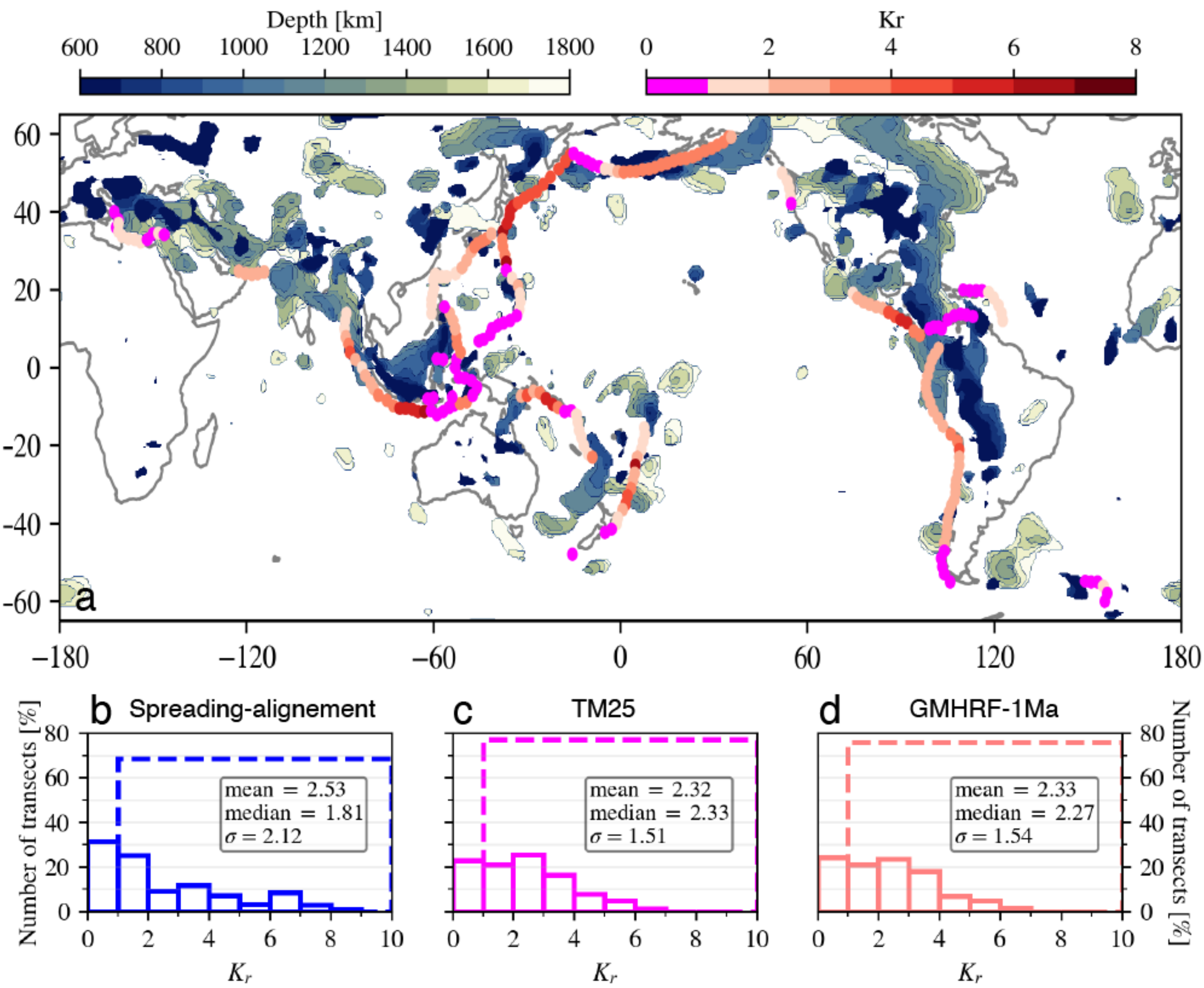


Figure 4: a) Map shows estimates of the kinematic ratio $K_r$ (reddish color scale) for all subduction transects of the SUBMAP database Heuret & Lallemand (2005). The regime of inferred slab folding ($K_r > 1$) prevails in most subduction zones. The absolute plate motion model is GMHRF-1Ma, which yields intermediate $K_r$ values compared to the also-investigated spreading-alignment and TM25 models (plotted in the supplementary information). Also shown in blue shades is the global inventory of subducted slabs in the lower mantle (600-1800 km depth). More precisely, these are contours of seismically fast P-velocity anomalies exceeding dvp/vp>0.35% in global model DETOX-P2 Hosseini et al. (2020). b) Histogram of $K_r$ values in the spreading-alignment reference frame. c) Histogram of $K_r$ values in the TM25 reference frame. d) Histogram of $K_r$ values in the GMHRF-1Ma reference frame.

Thus three separate lines of observational reasoning suggest that most present and past subduction zones feature(d) a surplus of subducted material not accommodated by lateral trench migration and slab sinking, so that instead slab folding is required. As discussed, common models of subduction dynamics (hereafter referred to as standard models) seldom reproduce this regime. Next we investigate whether adding a WAL to a standard model can shift its regimes from non-folding to folding over a wide range of model parameters.

# 3 Modeling approach

We use 2-D thermo-mechanical models of subduction dynamics. The governing equations are those suitable for multi-material, incompressible viscous flow, under the Boussinesq approximation, which are solved using the finite-element, control-volume, unstructured adaptive mesh Fluidity computational modelling framework, which has been carefully validated for simulations of this nature Davies et al. (2011); Kramer et al. (2012); Le Voci et al. (2014); Kramer et al. (2021). Our model setup and material properties are similar to Garel et al. (2014), albeit that in some cases we extend the models by incorporating a sub-lithospheric weak asthonospheric layer (WAL), similar to that in Suchoy et al. (2021). Below we summarize our modeling approach.

## 3.1 Model Setup

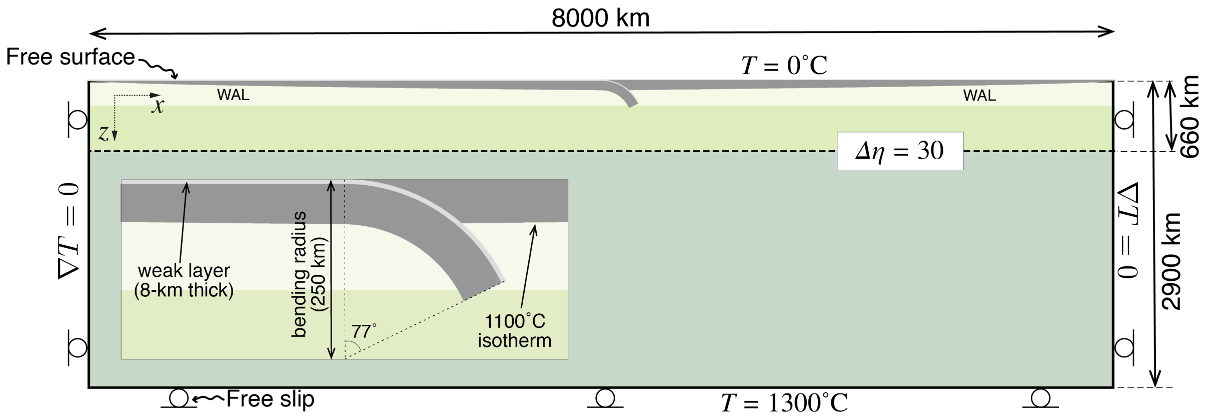


Figure 5: Model setup and boundary conditions. In the standard models the viscosity of the WAL is equal to that of ambient mantle ($\alpha = 1$), i.e., no weak asthenospheric layer (WAL) is present. We implement a viscosity contrast $\Delta\eta$ of 30 at 660-km depth, the boundary between upper and lower mantle. The initial curved geometry of the subducting plate is prescribed using a bending radius of 250 km, including the weak layer.

The model predicts the evolution of an isolated subduction zone comprising both a subducting plate (SP) and an overriding plate (OP), with no external forces or velocities applied to the system. The model domain is a Cartesian box that is 8000-km wide and 2900-km in height (i.e. the whole mantle depth). Mechanical boundary conditions on the sides and base of the domain are free-slip, with a free-surface at the top. We use no-flux thermal boundary conditions on the sides and impose constant temperatures of 0°C and 1300°C at the surface and at the bottom boundaries, respectively.

The initial temperature field is given by a half-space cooling model where the age of the plates vary linearly from the 0 at the ridges to ($A_{SP}$) for the subducting plate and to ($A_{OP}$) for the overriding plate. Models begin with a curved subducting slab to initiate subduction (see inset Fig. 5). Below the plates, the initial mantle temperature is equal to that of the bottom boundary.

We consider a composite visco-plastic rheology that accounts for four deformation mechanisms: linear diffusion creep, and non-linear dislocation creep, Peierls creep and pseudo-brittle yielding. The effective viscosity is:

$$\frac{1}{\eta_{eff}} = \left( \frac{1}{\eta_{diff}} + \frac{1}{\eta_{disl}} + \frac{1}{\eta_P} + \frac{1}{\eta_Y} \right) \tag{2}$$

which is bounded at lower and upper limits.

The diffusion ($\eta_{\text{diff}}$), dislocation ($\eta_{\text{disl}}$) and Peierls ($\eta_{\text{P}}$) viscosities follow the generic form:

$$\eta_{\text{diff|disl|P}} = A^{\frac{1}{n}} \exp\left(\frac{E+PV}{nRT_r}\right) \dot{\epsilon}_{II}^{\frac{1-n}{n}} \tag{3}$$

where $A$ is a prefactor, $n$ is the stress exponent, E and V are the activation energy and volume, respectively. $P$ is the lithostatic pressure, $R$ the gas constant, and $\dot{\epsilon}_{II}$ the second invariant of the strain-rate tensor. $T_r$ is the sum of model temperature and an adiabatic temperature gradient of 0.5 °C/km and of 0.3 °C/km in the upper and lower mantle, respectively. The pseudo-brittle yielding viscosity follows a yield-stress law

$$\eta_Y = \frac{\tau_Y}{2\dot{\epsilon}_{II}} \tag{4}$$

where the yield strength $\tau_Y = \min(\tau_0 + f_c P, \tau_Y^{\max})$, with $\tau_0$ the surface yield strength, $f_c$ the friction coefficient, $P$ the lithostatic pressure, and $\tau_Y^{\max}$ the maximum yield strength. The weak layer is 8-km thick, with a friction coefficient 10 times lower than the mantle material, and a maximum prescribed viscosity of $10^{20}$ Pa s. All rheological parameters are as in Table 1 of Garel et al. (2014) and we use a consistent solution strategy.

## 3.2 Treatment of WAL

The depth extent of a potential WAL is not well constrained. Some studies, which consider it to be a layer of partial melt, suggest that it is only 10-20 km thick Schmerr (2012); Sakamaki et al. (2013); Stern et al. (2015), whereas others advocate for a layer extending from the lithosphere-asthenosphere boundary up to 200-300 km depth (thus a thickness of approximately 100-200 km) Kawakatsu et al. (2009); Paulson & Richards (2009); French et al. (2013); Becker (2017); Barruol et al. (2019); Debayle et al. (2020). Here, we simulate the presence of the WAL by imposing a viscosity reduction between the 1100 °C isotherm (a proxy for the LAB) and a depth of 220 km, similarly to Suchoy et al. (2021). Note that models tested with a WAL extending up to 300 km depth showed little differences with the results reported below. We define the effective viscosity within the WAL as:

$$\eta_{\text{WAL}} = \alpha \eta_{\text{eff}} \tag{5}$$

where $0 < \alpha \leq 1.0$ is a reduction-viscosity factor.

The viscosity reduction of a WAL, and its origin, is also debated. For example, partial melt can lead to a 20-fold or larger viscosity reduction Holtzman (2016), but strongly depends on melt fraction, creep regime, grain size and wetting angle Kohlstedt & Zimmerman (1996). Milder viscosity reduction (< 5-fold) are expected from crystal-preferred orientation considerations Meyers & Kohlstedt (2021). We choose here to explore values of $\alpha$ in [0.2; 0.5; 1.0] (respectively, a viscosity reduction of 5-, 2-, or zero-fold – see below Section 4.1). Note that larger viscosity reductions ($\alpha$ of 0.01-0.1) have usually been used in global mantle flow models reproducing sub-plate seismic anisotropy Conrad & Behn (2010); Becker (2017). However, values of $\alpha$ lower than 0.1 in our set-up led to unrealistically large asthenospheric flow velocities (>50 cm/yr) and to thermal instabilities potentially associated to small-scale convection (further discussion in Section 5.3.1).

# 4 Model results

We first perform a set of simulations without a WAL ($\alpha = 1$), that we hereafter refer to as standard cases. Next, we explore sets of simulations with different degrees of weakening in the WAL (i.e. various values of $\alpha$), that we refer to as WAL cases. For each case, we define a reference simulation (with plate ages $A_{sp} = 40$ My and $A_{op} = 20$ My). We subsequently run simulations that span a range of initial ages to cover a wide range of strength and buoyancy for both plates, while being representative of all regions of the subduction regime diagram presented in Garel et al. (2014).

## 4.1 Standard cases - no WAL

### 4.1.1 Reference simulation [$A_{sp} = 40$ My;$A_{op} = 20$ My]

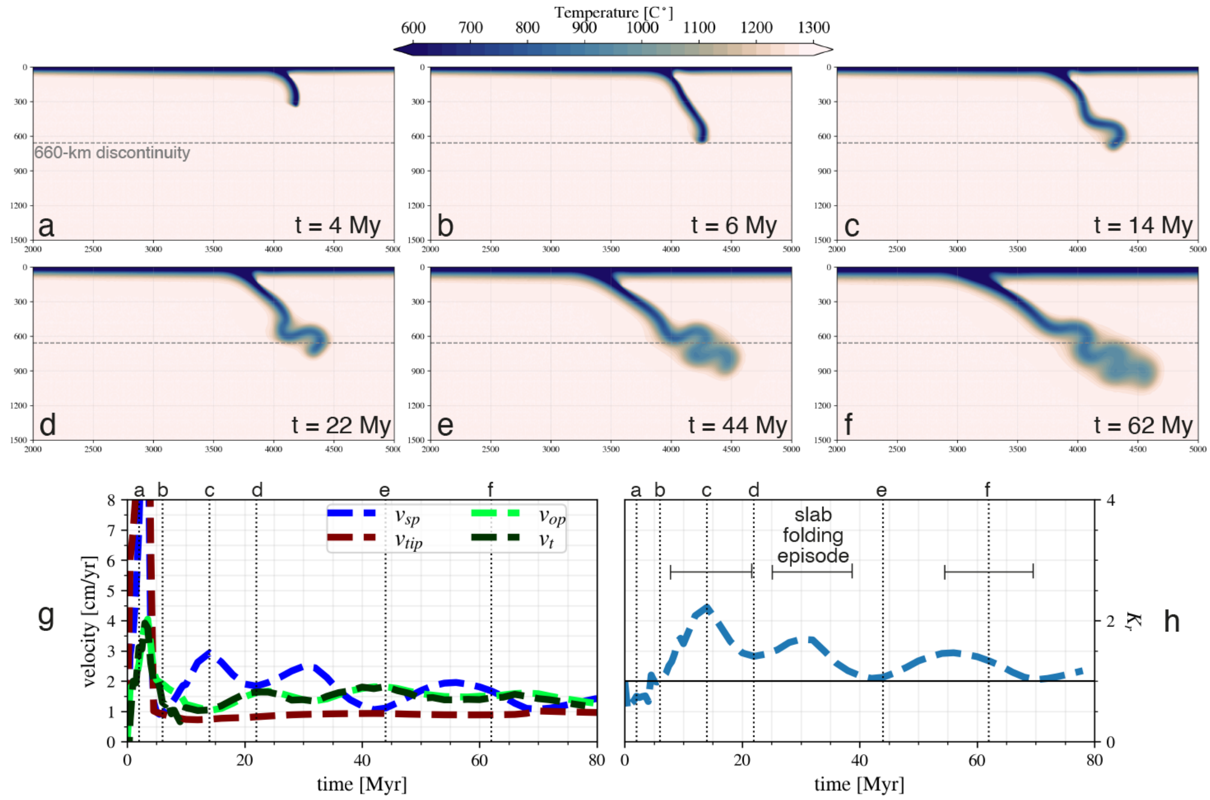


Figure 6: Reference standard model, featuring $A_{sp} = 40$ and $A_{op} = 20$ My. (a-f) Temporal snapshots of the evolution of the temperature field, from 4 to 62 My. g) Kinematics of plates, slab, and trench. $v_{sp}$ in dark blue, positive to the right. $v_{op}$ in green, positive to the left. $v_t$ in light blue, positive to the left. $v_{\mathrm{tip}}$ in dark red is the magnitude of slab tip velocity in. Vertical dashed lines mark snapshots times of (a-f). h) Temporal evolution of Kinematic ratio $K_r$, a proxy for the slab's propensity to fold (definition in the text). Horizontal black line denotes the boundary between the folding ($K_r > 1$) versus non-folding ($K < 1$) regime.

Figure 6 displays the temporal evolution of the reference simulation for the standard cases ($\alpha = 1$). During the first stage of free-sinking through the upper mantle (Fig. 6a), the subducting plate accelerates as slab pull increases with increasing slab length, reaching a peak velocity of $\sim$ 13 cm/yr with trench retreat/OP velocity peaking at $\sim$ 4 cm/yr (Fig. 6g). Since trench velocity is very similar to OP velocity, indicating little back-arc deformation, we will only describe the evolution in terms of trench velocity. As soon as the slab tip encounters the high-viscosity lower mantle at $t \simeq 5$ My, the plate velocities decrease to between 1 and 3 cm/yr. This slab-660 interaction is followed by two subsequent episodes of slab folding. First, the slab bends with an OP-wards concavity between $\sim$6 and $\sim$14 My (Fig. 6b-c), with an increase in SP velocity (up to 3 cm/yr) and low trench velocity of approximately 1 cm/yr. Then, a transient stage of slab rollback, associated with a slight increase in the trench velocities up to 1.5 cm/yr, lasts for approximately 5 My (Fig. 6c) and lowers slab dip in the upper mantle. The second folding episode occurs between $\sim$25 and $\sim$40 My: the deeper, folded portion of the slab flattens above the lower mantle while the shallow slab continues to roll back (Fig. 6d), increasing slab pull ($v_{sp}$ up to $\sim$ 2.5 cm/yr). Another transient stage of trench retreat without buckling follows (Fig. 6e), with trench velocity (1.8 cm/yr) greater than the SP velocity (1 cm/yr). From 50 My, a third slab-folding episode occurs, but with a smaller amplitude due to the obliquity of slab relative to the viscosity jump (Fig. 6f), and a smaller increase in $v_{sp}$. Overall, through time, all velocities decrease and tend towards 1 cm/yr, comparable to the sinking velocity of the deepest part of the slab within the lower mantle. Slab sinking rates in the lower mantle therefore strongly modulate, and perhaps even limit, surface kinematics.

The kinematic ratio, $K_r$, given in Equation 1 provides an alternative quantitative diagnostic. During the free-sinking stage, the slab-tip velocity reaches a peak value of 20 cm/yr, higher than the peak SP velocity $v_{sp}$ (13 cm/yr). As a consequence, the kinematic ratio $K_r \leq 1$ (Fig. 6g). After the slab has interacted with the 660-km discontinuity $K_r$ display oscillations. These oscillation follow those observed for $v_{sp}$ and $v_t$, when one of the two increases while the other decreases. Folding episodes occur when $K_r > 1$, while slab-retreating stages occurs for lower $K_r \simeq 1$. Through time, the amplitude of $K_r$ oscillations decrease, reflecting the decrease in folding as the slab inclines and the impact angle with the 660-km viscosity discontinuity decreases.

### 4.1.2 Slab morphologies and kinematic ratios across all standard cases

We run a series of no-WAL simulations with various initial plate ages (20-100 for the overriding plate; 10-100 for the subducting plate). Since we focus on the long-term evolution of these systems (i.e. well after the first stage of slab-free sinking through the upper mantle), Figure 7a only displays their state at $t = 80$ My.

Several studies have focused on the interaction and passage of slabs through the mantle transition and the resulting slab morphologies Torii & Yoshioka (2007); Billen (2010); Lee & King (2011); Čížková & Bina (2013); Billen & Arredondo (2018), sometimes characterizing a range of so-called subduction regimes Garel et al. (2014); Agrusta et al. (2017); Li et al. (2019); Briaud et al. (2020). Here we focus on two features after initial slab-660 interaction: trench motion and the amount of slab folding. Thus we define three regimes: strong trench retreat without slab folding (SR), strong trench retreat with slab folding (SRwF), and a weak trench retreat with slab folding (WRwF) (7a). The strong-retreat modes are those for which the total displacement of the trench during the simulation amounts to an average rate higher than 1 cm/yr, and weak-retreat modes when it is $le1$cm/yr. Following, Garel et al. (2014), the results of simulations are reported as functions of initial SP and OP ages, with the former controlling slab buoyancy and resistance to bending, and the latter controlling the OP bending resistance opposing trench retreat.

Note that due to our focus on the long-term trench motion and the tendency and nature of slab folding, the subduction regimes outlined herein differ from those used in Garel et al. (2014).

The SR regime in the simulations of the standard case occurs for both relatively old SPs and relatively old OPs. The regime WRwF occurs only for very young SPs. The regime that lies in between, SRwF, occurs over the widest parameter space. In simulations with relatively young OPs, only the SRwF is observed. For extremely young cases ($A_{SP} = A_{OP} = 20$ Myr), subduction is rapidly terminated through slab detachment, because the low slab pull cannot initially overcome the resisting forces.

Figure 7b displays the evolution of $K_r$ for four selected standard simulations. These simulations display peak $K_r$ of 1.5-2.2, shortly after the first slab-660 interaction (time range 5 to 20 Myr). Simulations [$A_{sp}$=40 My ; $A_{op}$=20 My] (ref. simulation - SRwF) and [$A_{sp}$=40 My ; $A_{op}$=65 My] (SRwF) display oscillations of $K_r$ associated with slab folding. Simulations [$A_{sp}$=65 My ; $A_{op}$=65 My] (SR) and [$A_{sp}$=100 My ; $A_{op}$=65 My] (SR) display $K_r \sim 1$ at all times after initial slab-660 interaction. At later times, the value of $K_r$ tends to 1, associated a decrease of both $v_{sp}$ and $v_s$ (see Figure S6 of Supp. Inf.).

## 4.2 WAL cases

We next perform simulations with a WAL, that is simulations where we impose values of the weakening factor $\alpha < 1$ in the sub-lithospheric mantle.

### 4.2.1 Reference WAL simulation with $\alpha = 0.5$ ($A_{sp}$=40 My and $A_{op}$=20 My)

In the reference WAL simulation (Fig. 8), the first slab-660 interaction occurs at $\sim$ 1.5 My, earlier than in the comparable standard simulation ($\simeq$4 My). The first slab buckling episode occurs shortly after at 2-10 My (Fig. 8a-b), with subducting plate and trench velocities of 5.5 cm/yr and 1 cm/yr, respectively. A second folding episode takes place after 10 My (Fig. 8c-d) during which the SP velocity increases from 1.8 to 3.8 cm/yr between $t = 14$ My and $t = 22$ My and that of the trench decreases from 1.8 to 0.8 cm/yr. The next folding episode (between $t = 35$ My and $t = 55$ My, Fig. 8e) is associated with a stationary trench, while the SP velocity stabilizes at $\sim$ 2.5 cm/yr. A third fold forms after $t = 55$ My (Fig. 8f) which produces a peak subducting-plate velocity of 4.8 cm/yr.

The opposite evolution of trench and SP velocity, associated with $K_r$ oscillation, is even more apparent that in the reference simulation of the standard case. After the first slab-660 interaction, the slab tip velocity remains nearly constant at around 1 cm/yr, independent of slab folding and oscillation of surface velocities.

### 4.2.2 Mantle drag forces on the subducting plate

Differences in the evolution of reference simulations can be explained by a reduction of mantle drag at the base of the subducting plate. We calculate a drag force as the integral of the tangential stress along the 1100°C isotherm (in N m$^{-1}$). Figure 9 displays the temporal evolution of this diagnostic for the reference simulations (i. e, with and without a WAL).

Prior to initial slab-660 interaction, both models show sub-lithospheric mantle moving towards the trench but with velocities reduced relative to overlying lithosphere (Couette-flow type). Shear stresses beneath the subducting plate are positive and mostly negative beneath the overriding plate. Shear stresses along the base of the subducting plate remains

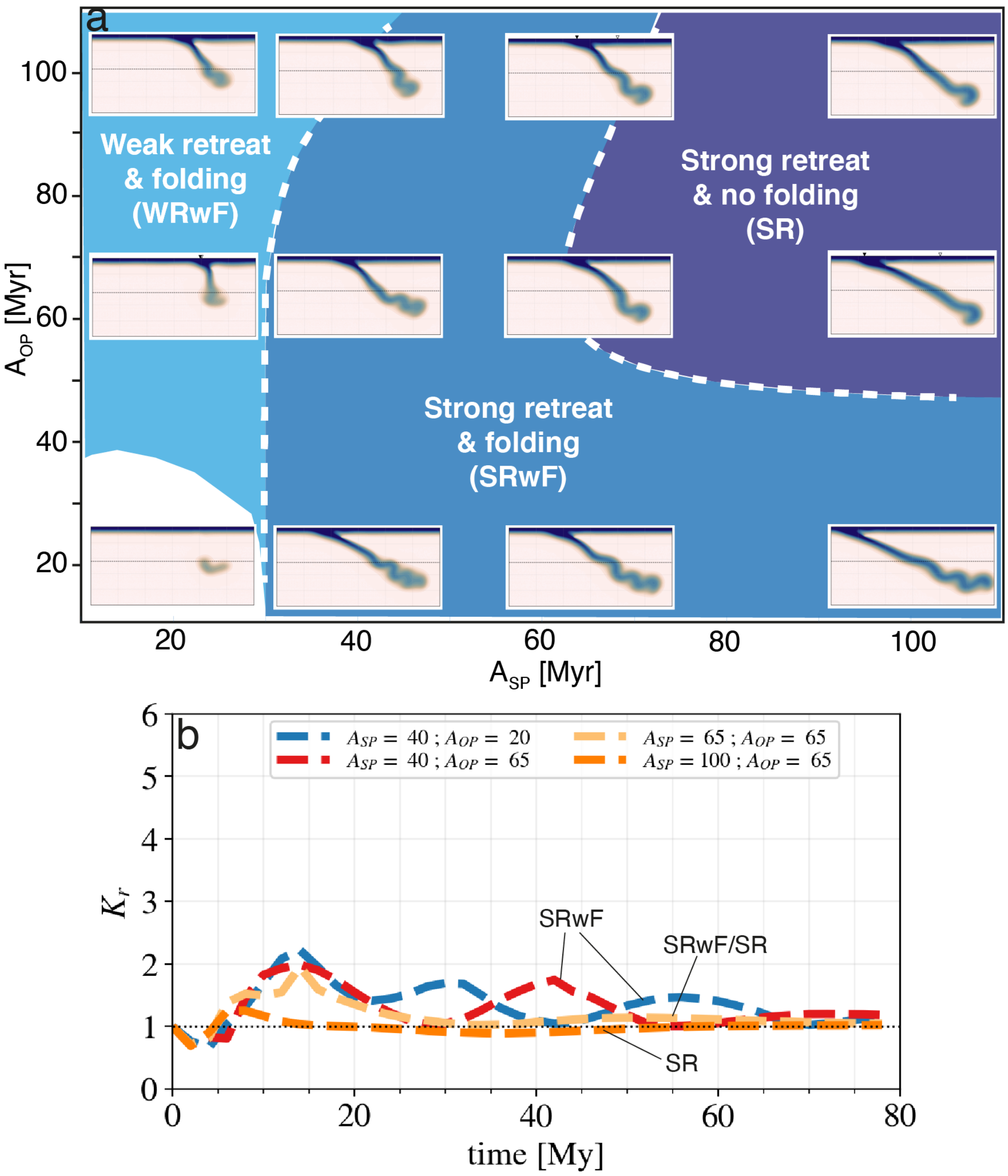


Figure 7: a) Snapshots of the final state (after model run times of 80 My) of all standard models superimposed on a regime diagram. The three regimes are: strong trench retreat without slab folding (SR, purple), strong trench retreat with slab folding (SRwF, dark blue), and weak trench retreat with slab folding (WRwF, light blue). The boundaries between regimes are approximate. b) Kinematic ratio $K_r$ as a function of the time since the initiation of subduction, for four of the standard models shown in (a). The subduction regimes associated with the evolution of those models is indicated by labels. Over time, all four models tend towards no-folding ($K_r \approx 1$).

positive after initial slab-660 interaction (Fig. 9c,e), and the drag force remains negative. Overall, we observe that the absolute value of the drag force beneath the subducting plate decreases with time as a consequence of the reduction in the length of that plate with time.

The two simulations display similar oscillatory trends that reflect slab folding behavior. However, as expected, the absolute drag is lower in the simulation featuring a WAL,

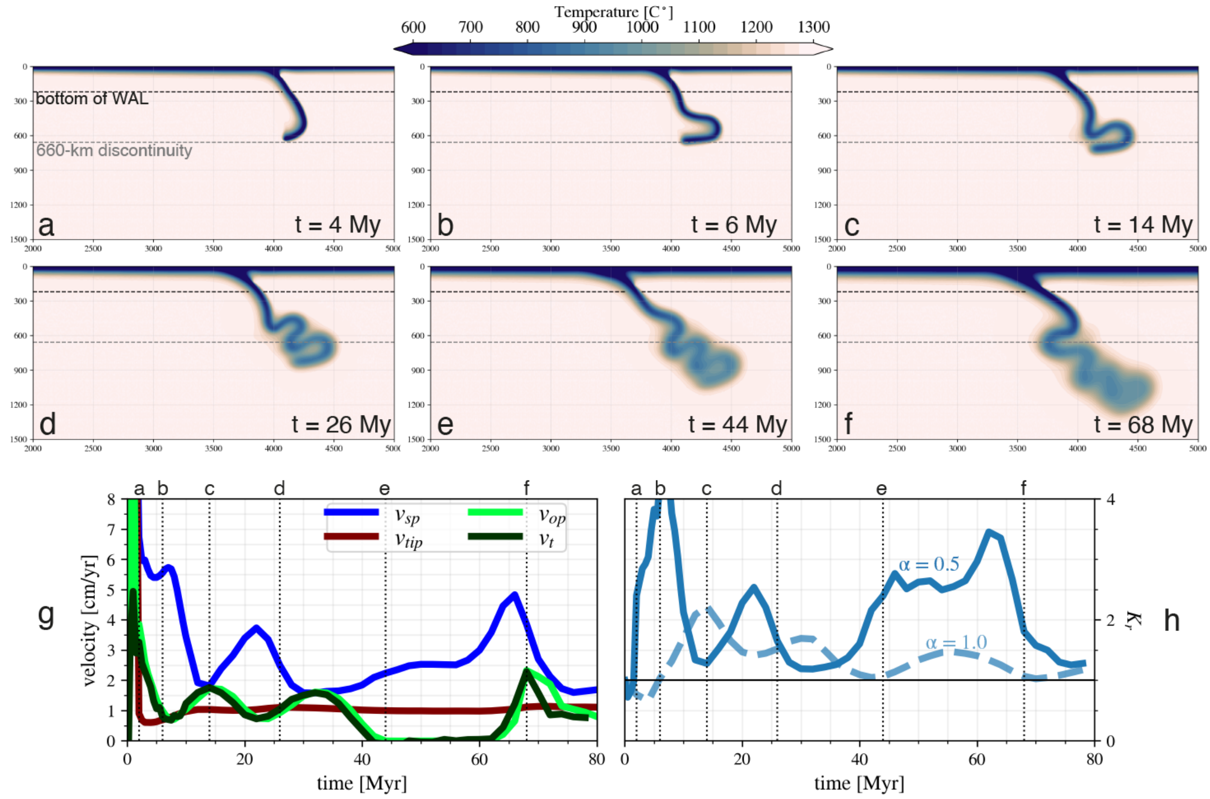


Figure 8: Reference simulation [$A_{sp}$=40 My ; $A_{op}$=20My] of the WAL case with a two-fold viscosity reduction ($\alpha = 0.5$). Panels and plotting styles as in Fig. 6. Panel (h) displays the evolution of $K_r$ for this reference WAL case (solid blue) and also for the reference standard case of Fig. 6 (dashed blue).

explaining the higher SP velocities observed in this case (Fig. 8g) relative to the comparable no-WAL case (Fig. 6g). The faster subducting plate may hamper trench retreat with faster trench-wards asthenospheric flow opposing slab rollback Alsaif et al. (2020). Slab sinking, and hence a faster subducting plate, may also be enhanced through dislocation creep decreasing the viscosity of adjacent asthenosphere, further facilitating slab descent Garel et al. (2020). All these effects favor the higher $K_r$ observed in the reference WAL case compared to the reference standard case.

#### 4.2.3 Subducting slab morphologies and kinematic ratios in the WAL case

As with the standard cases, we run a series of WAL simulations with $\alpha = 0.5$ spanning plate ages that range from 20-100 My. Figure 10a displays their final state at 80 My, together with the inferred regime diagram. Consistent with the standard cases, WAL cases exhibit three regimes (SR, SRwF, and WRwF), but regime boundaries are shifted towards higher plate ages. In particular, the WAL simulations [$A_{sp}$=65 My-$A_{op}$=65 My] and [$A_{sp}$=65 My-$A_{op}$=100 My] now lie more clearly in the SRwF regime while their standard equivalents belong to the SR regime. Moreover, the WAL simulations with $A_{sp} = 40$ My lie within, or very close to, the WRwF regime, whereas for standard cases, only those with $A_{sp} \geq 20$ My are within this regime.

Figure 10b displays the kinematic ratio $K_r$ of selected WAL simulations. As in the standard cases, the ratios $K_r$ before and during the first slab-660 interaction is generally

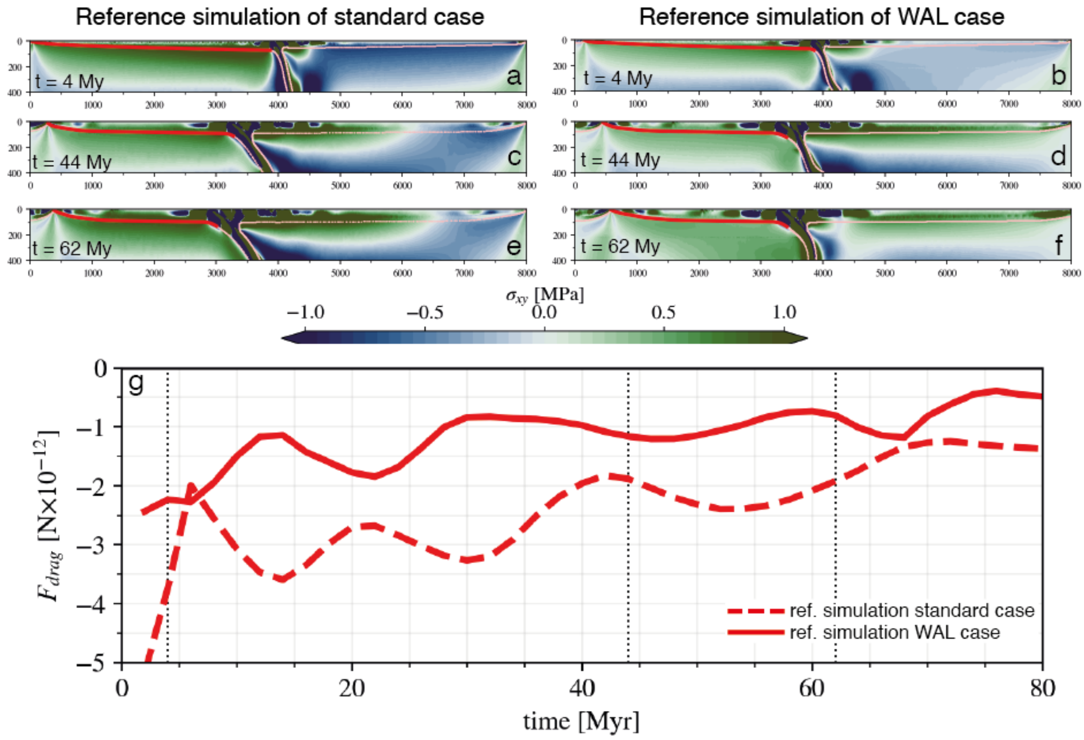


Figure 9: Viscous mantle resistance acting below the subducting plate in the reference standard model versus in the reference WAL model. (a,c,e) Snapshots of the shear stress field (background colour) the reference standard model at times 4 My, 44 My, and 62 My. (b,d,f) Same snapshots for the reference WAL model. Red line traces the LAB isotherm of 1100°C (pink) used to calculate the drag below the subducting plate. Positive shear stresses imply that the tangential component of the stress vector - calculated along the quasi-horizontal LAB isotherm of the subducting plate - is toward the left. g) Evolution of the mantle drag force below the subducting plate for the reference standard model (dashed red) and the reference WAL model (solid red) . Negative values denote a force toward the left. The total mantle drag force onto the subducting plate is negative in accordance with the stress vector. Three vertical dashed lines indicate the times of the snapshopts (a-f). The drag force beneath the overriding plate is less straightforward to analyze, see Supporting Information section S4.

higher than 1. Some peak values of $K_r$ reached in the WAL simulations are even greater ($>$ 2) than the highest values observed in the standard simulations (7b). Most importantly, two of these simulations [$A_{sp}$ = 40 My;$A_{op}$ = 20 My] (ref. simulation for the WAL case) and [$A_{sp}$ = 65 My;$A_{op}$ = 65 My] display $K_r > 2$ even after the first slab-660 interaction: the presence of a weak layer favors the excess accumulation of subducted material in the mantle relative to the accommodation by motion of both the trench and slab tip, resulting in substantial slab folding. The simulation [$A_{sp}$=100 My ; $A_{op}$=65 My] shows values of $K_r$ close to 1 at all times after initial slab-660 interaction, consistent with the standard case, lying in the SR regime. WAL simulation [$A_{sp}$=65 My ; $A_{op}$=65 My] exhibits intermediate behavior, with oscillations of $K_r$ up to 1.5, while its standard equivalent show values close to 1. This is because the former clearly lies in the SRwF regime while its equivalent standard case lies near the transition from the SRwF to the SR regime.

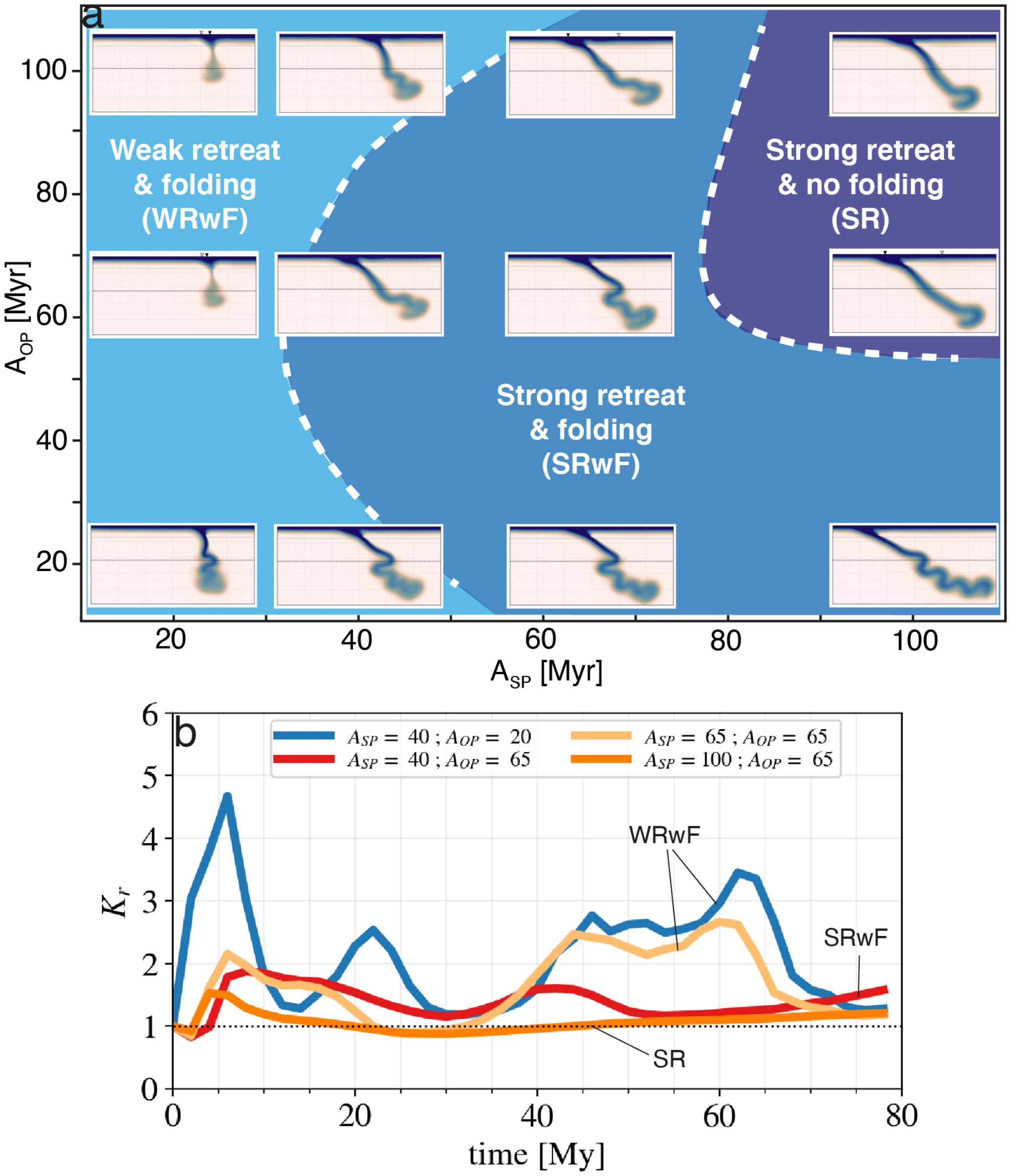


Figure 10: a) Regime diagram of WAL models with two-fold viscosity reduction ($\alpha = 0.5$). Panels and plotting styles as in Fig. 7.

### 4.2.4 WAL cases with $\alpha = 0.2$

We finally run a series of simulations for a WAL with $\alpha = 0.2$ (Fig. 11a). The most striking difference, relative to all previous cases examined, is that the SR regime no longer appears within the range of plate ages investigated: slab folding behavior is consistently observed, throughout the investigated parameter space. Moreover, the boundary between the WRwF and SRwF regimes shifts to values of subducting plate ages higher than 65 My and for overriding plate ages higher than 20 My. Hence, strong retreat ($v_t > 1$ cm/yr) now only occurs for heavy and stiff plates, with the WRwF regimes becoming dominant. In particular, when the overriding plate is relatively weak ($A_{op} \simeq 20$ My) vertical slab folding piles in the lower mantle are ubiquitous. Cases also exhibit higher $K_r$ values (1.5-4) well after the first slab-600 interaction (Fig. 11b).

Finally, it is worth noting that in the majority of the above simulations with a WAL and $\alpha = 0.2$, thermal instabilities form within the weak layer at the base of the plates. These transient instabilities, which take the form of drips of cold lithosphere generated beneath both overriding and subducting plates, develop as the lithosphere thickens in response to conductive cooling. The drips are then advected by lateral mantle flow and mix with underlying mantle. We note that they only occur when the asthenospheric viscosity reaches values close to or below $\sim 10^{19}$ Pa s, as advocated by previous studies van Hunen et al. (2003); Ballmer et al. (2011); Le Voci et al. (2014); Davies et al. (2016). Further analyses of these features will be the focus of a future study.

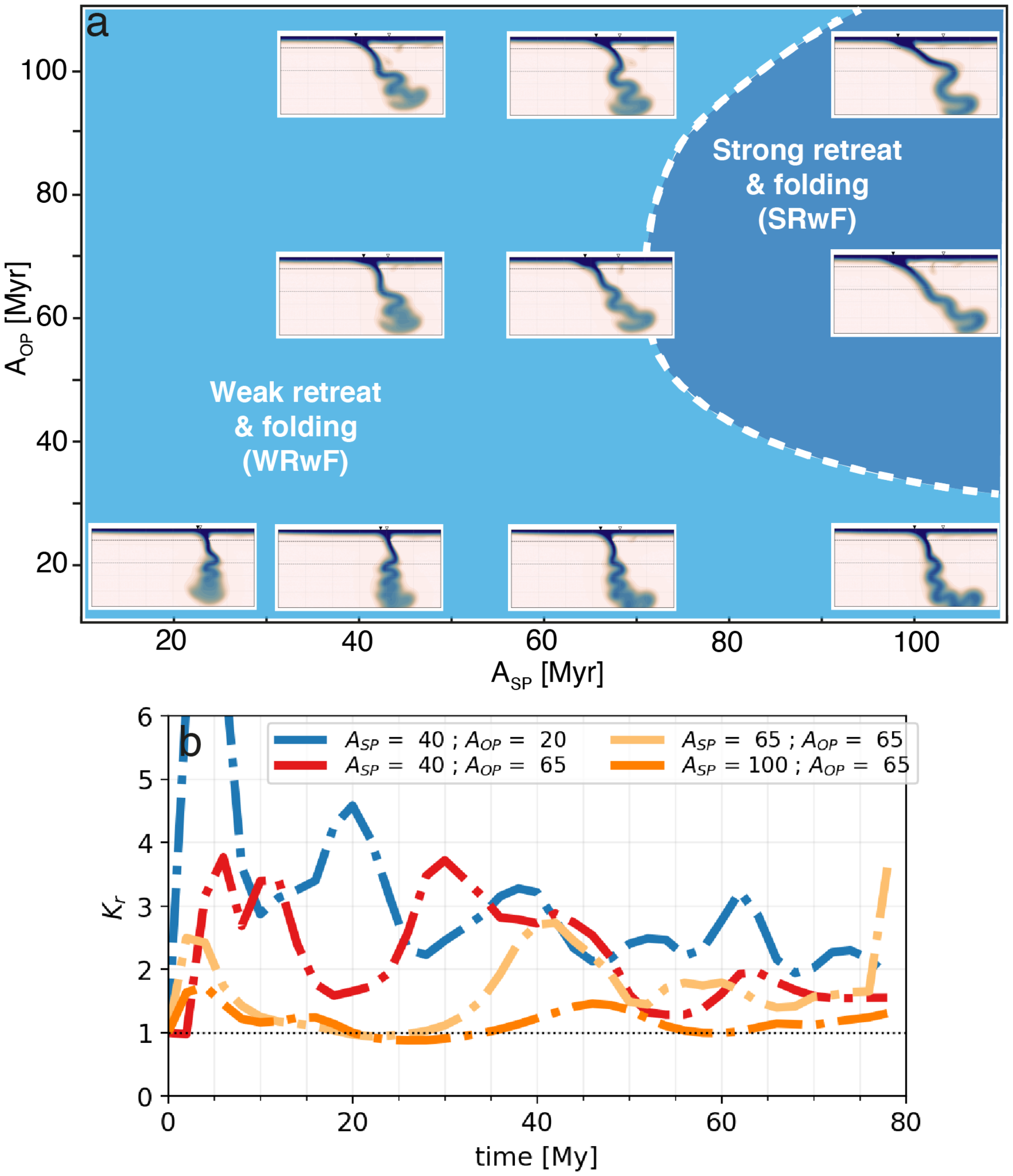


Figure 11: Regime diagram of WAL models with a five-fold viscosity reduction ($\alpha = 0.2$). Panels and plotting styles as in Fig. 7. The models remain in the weak retreat and folding regime over the entire run time, unless they feature very old plate ages.

# 5 Discussion

## 5.1 Surface velocities and kinematic ratios

On Earth, subduction rates are typically 3 to 5 times higher than absolute trench velocities, and 5 times higher than estimated slab sinking velocities (see Section 2). This yields kinematic ratio $K_r$ estimates above 1 for most subduction zones, creating a scenario that favours a slab thickening/folding regime (see section 2.2).

Simulations without a WAL produce surface kinematics at odds with these observational constraints because once the slab has interacted with the transition zone, the subducting plate $v_{sp}$ slows down to approach slab sinking rates of 1 cm/yr. The addition of a WAL renders the simulations more compatible with observational constraints, in that $v_{sp}$ up to 5 cm/yr are maintained long after initial slab-660 interaction (Fig. S4-S5 in Supp Info), and trench velocities $v_t$ are attenuated to typically lower than 1 cm/yr. Our models with WAL thus reproduce the rapid subduction rate, near-stationary trenches, and slow slab-sinking rates observed on Earth. WAL simulations have higher $K_r$ values than the standard models, as summarized by Figure 12. The time-averaged kinematic ratio $\bar{K}_r$ (after initial slab-660 interaction, i.e., averaged between 20-80 My) ranges from 1.0–1.3 in the standard models (Fig. 12a, except for the youngest SP plate ages). In contrast, $\bar{K}_r$ ranges between 1.0–3.1 in WAL simulations with two-fold weakening ($\alpha = 0.5$, Fig. 12b), and between 1.0–2.9 for five-fold weakening ($\alpha = 0.2$, Fig. 12c). $\bar{K}_r$ is generally higher in WAL simulations with $\alpha = 0.2$ (although the maximum value of $\bar{K}_r = 3.1$ occurs for $\alpha = 0.5$ and the youngest plate ages).

WAL simulations also show higher peak values of (non-averaged) $K_r$, before and after first slab 660-interaction. In the standard models, $K_r$ mostly ranges between 1–2 (see Fig. 7 and Supp Info Fig. S6a), whereas the WAL simulations exhibit peak $K_r$ values above 2 and up to 6-7 (Figs. 10 and Fig. 11 – see also Figs. S6b,c in Supp. Info.). Hence only the models with a weak layer produce kinematic ratios $K_r$ that are comparable to those estimated for subduction zones in nature (Fig. 4).

Behr & Becker (2018) have suggested the lubrication effect of a weak sedimentary layer above the subducting plate as an alternative mechanism for increasing $v_{sp}$ in models of subduction dynamics ?see also>[]duarte2013three. They showed that $v_{sp}$ could increase by one to two orders of magnitude if sediments reduced viscous resistance at the interface by a comparable amount. However, recent models have shown that a weaker plate interface also favors an increase in trench retreat $v_t$ Pusok et al. (PREPRINT); Behr et al. (2022). We investigated this by running a simulation without a WAL but featuring a plate interface layer with a two-fold reduction in maximum viscosity (see Fig. S8). Relative to our comparable standard case, both $v_{sp}$ and $v_t$ increased slightly, resulting in a negligible difference to $K_r$. This suggests that a weaker plate interface would not allow us to reconcile model predictions with the available observational constraints on $K_r$, further supporting an important role for a WAL.

## 5.2 Slab morphologies

The presence of a WAL strongly impacts the subduction regimes and lower mantle slab morphologies, as encapsulated by the proxy of $K_r$. Simulations without a WAL produce low-to-moderate values of $K_r$, and moderate-to-high trench retreat rates. Without a WAL, strong-retreat regimes are thus dominant across the parameter space examined, and only models with the youngest, weakest overriding plate (20 My) exhibit some slab-folding

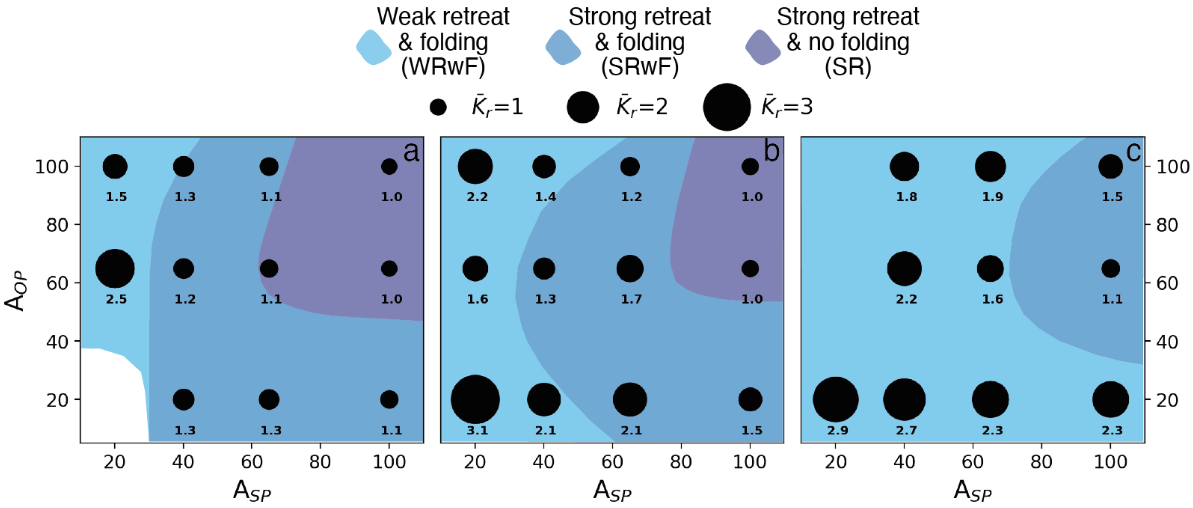


Figure 12: Summary of modeling runs and regime diagrams. Time-averaged kinematic ratio $\bar{K}_r$ after the first slab-600 interaction, plotted in the two-dimensional space of SP and OP plate ages (unit Myr). for (a) standard non-WAL models; (b) WAL models with $\alpha = 0.5$; (c) WAL models with $\alpha = 0.2$. $\bar{K}_r$ is represented by black dots with sizes proportional to $\bar{K}_r$ values, which are also printed. Background colors denote folding regimes as in previous figures. The time-averaging window for obtaining $\bar{K}_r$ is [20-80] My.

behaviour. However, even for a WAL with a moderate viscosity decrease ($\alpha = 0.2$), the (non-folding) SR regime disappears and the SRwF regime only occurs in simulations with relatively stiff and buoyant plates ($A_{sp} > 80$ My,$A_{sp} > 40$ My). Hence, Slabs folding and vertically piling in the lower mantle, beneath near-stationary trenches, become the prevailing morphologies of the WAL simulations (light blue shading in Fig.12). These results demonstrate, for the first time, that models of subduction dynamics (without external forcing) are able to produce lower-mantle slab morphologies observed by tomography, while also honouring the plate and trench velocities measured at the surface.

We note that the amplitude of lower-mantle slab folds in our simulations is consistent with theoretical predictions based on a thin-sheet mathematical formulation. Ribe (2003) and Ribe et al. (2007) used these formulations to derive a scaling law for the amplitude of folds of a vertically descending, viscous sheet that buckles as it encounters resistance at a sharp viscosity jump, or a rigid barrier. The predicted fold amplitude is half the fall height, which would be half the thickness of the upper mantle in the context of subduction: approximately 330 km. Our simulations with more pronounced vertical slab folding produce 300 to 500-km wide folds in the lower mantle, that are consistent with this theory. We note that the presence of a WAL enhances the frequency of folding in the models but leaves their width reasonably unchanged. The modeled fold amplitudes of 300-500 km are moderately smaller than the 400-700 km wide "slab walls" imaged by seismic tomography (e. g. Sigloch & Mihalynuk, 2013). It remains to be investigated whether this difference is due to shortcomings of the physical approximations used in our dynamic models, or due to tomographic blur.

From models of subduction dynamics, it has been suggested that sustained, quasi-periodic slab folding, over tens of millions of years after initial slab-660 interaction, can occur only if the mineralogical phase transition around 410 and 660 km were included in the models Běhounková & Čížková (2008); Čížková & Bina (2013); Agrusta et al. (2017); Briaud et al. (2020), and/or if the subducting slab was quite weak, e.g., made of young seafloor Garel et al. (2014); Agrusta et al. (2017); Strak & Schellart (2021). While we acknowledge that these factors may further enhance slab folding, we stress that our simulations with a WAL did not require the phase transitions in order to produce sustained slab folding. The Clapeyron slopes of the phase transitions remain under discussion ?see e.g.>[and references therein]agrusta2017subducting, so their relative role in slab folding remains to be clarified. In a similar vein, the inclusion of a WAL yielded slab folding of relatively thick and stiff subducting plates (Fig.12). No additional slab-weakening mechanism or slab-buoyancy variation was required. We note that it has also been suggested that vertical piles of lower-mantle slabs could only be produced in the context of a fixed overriding plate Lee & King (2011); Běhounková & Čížková (2008); Čížková & Bina (2013); Billen & Arredondo (2018). Here we have demonstrated that vertical slab folding slab can also occur in simulations with a WAL, in which trench retreat remains self-consistently limited (Fig. 8).

# Conclusion

Previous numerical and analogue models of subduction dynamics tend to produce surface kinematics and lower-mantle slab morphologies that do not match first-order observational constraints. We have shown that including a weak asthenospheric layer below the lithosphere into numerical models of subduction dynamics eliminates these mismatches. The lubricating effect of the asthenosphere produces a velocity increase of the subducting plate and a reduction of trench retreat, yielding predicted velocities that closely match those recorded on Earth. These velocity changes are sustained long after the subducting slab has penetrated into the lower mantle. The surplus of rapidly subducting lithosphere is accommodated by folding, rather than by accelerating trench retreat or slab sinking. This leads to an apparent horizontal widening of the slab in the lower mantle, as is observed by seismic tomography. Substantial near-vertical slab piles accumulate over time because trench motion is limited. We find that a viscosity reduction below the plate by a factor of only 2 to 5 is sufficient to completely shift the dynamics in these models – from non-folding with slow subduction and substantial trench retreat, to regimes of multiply folded, wall-like slab piles under near-stationary trenches. The latter then dominate across a wide parameter space of subducting and overriding plate ages. Our results provide strong independent support for the presence of a weak asthenospheric layer beneath Earth's lithosphere.

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