

1           **Efficient Probabilistic Prediction and Uncertainty Quantification of**  
2                                   **Hurricane Surge and Inundation**

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10 ABSTRACT: This study proposes a methodology for efficient probabilistic prediction of near-  
11 landfall hurricane-driven storm surge, tide, and inundation. We perturb forecasts of hurricane track,  
12 intensity, and size according to quasi-random low-discrepancy Korobov sequences of historical  
13 forecast errors with assumed Gaussian and uniform statistical distributions. These perturbations  
14 are run in an ensemble of hydrodynamic storm tide model simulations, and the resulting set of  
15 maximum water surface elevations are used as a training set to develop a Polynomial Chaos (PC)  
16 surrogate model from which global sensitivities and probabilistic predictions can be extracted. The  
17 maximum water surface elevation is extrapolated over dry points incorporating energy head loss  
18 with distance to properly train the surrogate for predicting inundation. We find that the surrogate  
19 constructed with 3rd order PCs using Elastic Net penalized regression with Leave-One-Out cross-  
20 validation provides the most robust fit across training and validation sets. Probabilistic predictions  
21 of maximum water surface elevation and inundation area by the surrogate at 48-hour lead time  
22 for three past U.S. landfalling hurricanes (Irma 2017, Florence 2018, and Laura 2020) are found  
23 to be reliable when compared to best-track hindcast simulation results, even when trained with  
24 as few as 19 samples. The maximum water surface elevation is most sensitive to perpendicular  
25 track-offset errors for all three storms. Laura is also highly sensitive to storm size and has the least  
26 reliable prediction. This methodology is built into an open-source Python framework available  
27 from <https://github.com/noaa-ocs-modeling/EnsemblePerturbation>.

28 SIGNIFICANCE STATEMENT: This purpose of this study is develop and evaluate a methodol-  
29 ogy that can be used to provide high-quality probabilistic predictions of hurricane-induced storm  
30 surge and inundation with limited time and resources. This is important for emergency management  
31 purposes during or after the landfall of hurricanes. Our results show that sampling forecast errors  
32 using quasi-random sequences combined with machine learning techniques that fit polynomial  
33 functions to the data are well-suited to this task. The polynomial functions also have the benefit  
34 of producing exact sensitivity indices of storm surge and inundation to the forecasted hurricane  
35 properties such as path, intensity, and size, which can be used for uncertainty estimation. The code  
36 implementing the presented methodology is publicly available on Github.

## 37 **1. Introduction**

38 Tropical and subtropical storms build up storm surges that affect populated coastal regions in  
39 the U.S. and internationally. The temporarily-higher sea levels from these storm surges result in  
40 widespread inundation of coastal low-lying areas, invoking flood and wave damage to residential  
41 and commercial structures. Storm surges from named storm events are estimated to cause billions  
42 of dollars in damages in the U.S. annually (NOAA National Centers for Environmental Information  
43 (NCEI) 2022). Under the requirements of the part of the Consumer Option for an Alternative  
44 System to Allocate Losses (COASTAL) Act, the National Oceanic and Atmospheric Administration  
45 (NOAA) is responsible for determining the extent of storm surge and storm tide to inform response  
46 and application of relief funding from the Federal Emergency Management Administration (FEMA)  
47 after a storm event.

48 While currently not in operation, in this project we are investigating the application of a Hurricane  
49 Surge On-demand Forecast System (HSOFS) that could be employed when a tropical cyclone (TC)  
50 approaches and makes landfall along U.S. coastlines to provide predictions of hurricane-driven  
51 storm surge and inundation (Vinogradov et al. 2018). HSOFS uses a hydrodynamic storm tide  
52 model to simulate coastal water levels and inundation on high-resolution unstructured meshes,  
53 which may also be coupled to a wind-wave model to capture wave setup effects (Dietrich et al.  
54 2011; Moghimi et al. 2020). The system would be utilized to produce either; 1) near-landfall  
55 forecasts for support of recovery and response in the immediate aftermath of hurricane landfall, or  
56 2) hindcasts for allocating flooding-related insurance losses as part of the COASTAL Act (Abdolali

57 et al. 2021). But as with any modeling, the uncertainty in the results are dependent on the uncertainty  
58 and accuracy of the input parameters, predominantly those of the hurricane track, intensity, and  
59 size. Therefore, it is becoming increasingly important to provide probabilistic predictions and  
60 uncertainty estimates for decision making. Obtaining the probabilistic result makes the predictions  
61 more informative and robust, and reduces the likelihood of overestimation or underestimation of  
62 the severity of storm surge.

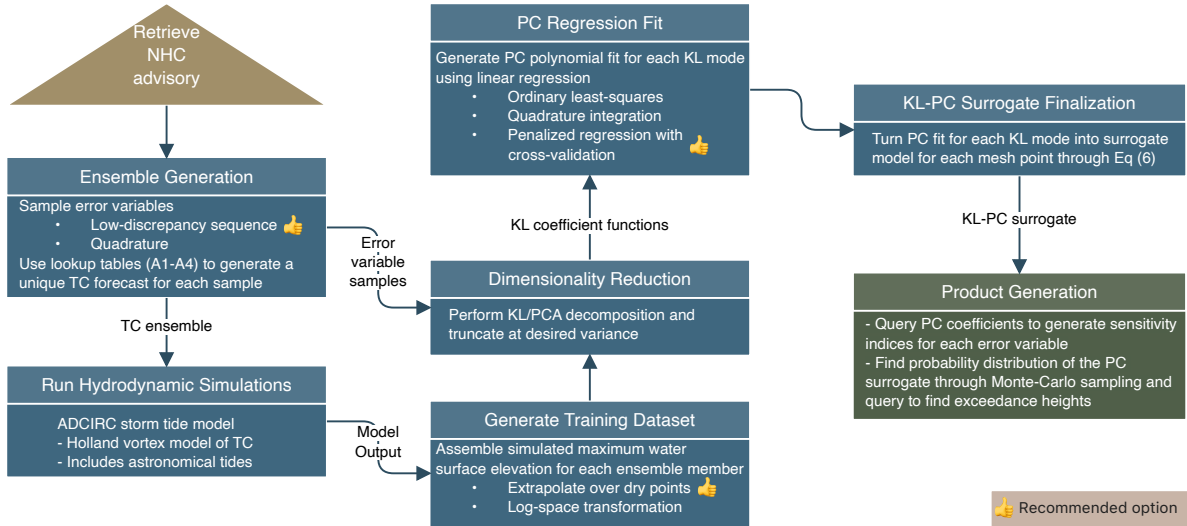
63 However, the complexity of HSOFS, which was previously in operation, leads to a relatively high  
64 computational load, limiting the number of model ensembles achievable in a time and resource-  
65 limited environment. This has been an obstacle in the development of a probabilistic version.  
66 In contrast, the National Hurricane Center’s (NHC) Probabilistic Tropical Storm Surge (P-Surge)  
67 model (Taylor and Glahn 2008) performs hundreds of ensemble simulations within the allotted one  
68 hour time-frame (~30 min per simulation per CPU) through an ad-hoc full factorial perturbation  
69 of estimated hurricane track, intensity and size errors. P-Surge is based on the Sea, Lake, and  
70 Overland Surges from Hurricanes (SLOSH) hydrodynamic code (Jelesnianski et al. 1992), which  
71 makes several physical simplifications (Joyce et al. 2019) and employs limited-area meshes for  
72 efficiency. Thus, the aim of this study is to develop an efficient ensemble prediction framework  
73 (requiring few model simulations) that can be used by the more comprehensive and computationally  
74 intensive HSOFS model for accurate near-landfall probabilistic forecasts of hurricane surge and  
75 inundation.

76 To this end, Davis et al. (2010) divided the range of the hurricane track errors into equal-area bins  
77 depending on a user-defined priority level, and estimated that 27 ensemble members resolved 90%  
78 of inundation. Additionally, Kyrioti et al. (2021a) showed that quasi-Monte Carlo methodologies  
79 can be used to improve sampling efficiency of hurricane parameter errors over the full factorial  
80 approach used by P-Surge. However, there may still be limitations in the information available  
81 from smaller model ensembles [ $O(10)$ ] that we aim for here. A possible solution is to seek a  
82 surrogate approximation that can be used to rapidly sample a wider distribution of input hurricane  
83 parameters and obtain robust statistical quantities, without having to query and iterate over the  
84 costly hydrodynamic model.

85 Many such surrogate models for storm surge prediction have been proposed, using machine  
86 learning techniques such as Gaussian Processes (GP; kriging), artificial neural networks (ANN),

87 and convolution neural networks (CNN), often combined with dimensionality reduction and k-  
88 means clustering via Principal Component Analysis (PCA) (e.g., Jia and Taflanidis 2013; Taflanidis  
89 et al. 2013; Kim et al. 2015; Hashemi et al. 2016; Lee et al. 2021; Kyprioti et al. 2021b; Plumlee  
90 et al. 2021). This approach often involves training a surrogate model using a large ensemble  
91 of synthetic hurricanes which can then be used to predict the storm surge based on the current  
92 hurricane parameters as inputs (Taflanidis et al. 2013; Kim et al. 2015; Hashemi et al. 2016; Lee  
93 et al. 2021). As noted by Lee et al. (2021), one of the limitations of this approach is that nonlinear  
94 interactions of surge with other processes (e.g., astronomical tides, background sea levels, and  
95 rainfall) are ignored, which could be particularly important for inundation behavior. A potential  
96 solution is to generate a new surrogate model for the current storm that includes (some of) these  
97 interactions in the hydrodynamic model, as most recently explored by Plumlee et al. (2021) using  
98 GPs.

99 In this study we also seek a solution that develops a surrogate model on-the-fly to provide both  
100 robust statistics and uncertainty information of storm surge and flooding predictions for the current  
101 storm. A method potentially well-suited to this application is Polynomial Chaos (PC) theory,  
102 which has been recently used for developing probabilistic predictions and analyzing the sensitivity  
103 of surge to hurricane parameters with good success (Sochala et al. 2020; Ayyad et al. 2021). PC  
104 is a convenient means to propagate uncertainties from inputs to outputs of interest for general  
105 computational models (Sargsyan 2017). It can further be interrogated to rapidly evaluate moments  
106 and sensitivities due to their analytical availability, or quantiles and probability density functions  
107 (PDFs) via computationally inexpensive sampling. Thus, in this study we adopt PC theory and  
108 develop strategies around efficient random variable sampling, dimensionality reduction, penalized  
109 regression with cross-validation, and manipulation of the training set to optimize the setup for PC  
110 construction. We evaluate the accuracy of this PC-based surrogate model and demonstrate the  
111 reliability of the probabilistic prediction for three historical U.S. landfalling hurricanes (Irma 2017,  
112 Florence 2018, and Laura 2020). Statistical quantities and variance-based sensitivities from the  
113 PC surrogate can be distributed along with surrogate itself as a product of the ensemble HSOFS  
114 modeling system. The ensemble generation and PC analysis methodology presented in this paper  
115 is implemented in an open-source Python framework called `EnsemblePerturbation`.



121 FIG. 1. Flowchart of the proposed methodology for efficient probabilistic predictions and uncertainty quantifi-  
 122 cation of hurricane storm surge and inundation.

## 116 2. Methods and Experiment

117 A flowchart of the proposed methodology in this study is shown in Fig. 1. To fully comprehend  
 118 components of the flowchart we refer the reader to the rest of this section (a-d), as well as to the  
 119 results section 3 for details. Finally, in section e we describe the experiments we conduct to assess  
 120 the accuracy of different options and evaluate the reliability of the probabilistic prediction.

### 123 a. Storm Surge Modeling Component of the On-demand System

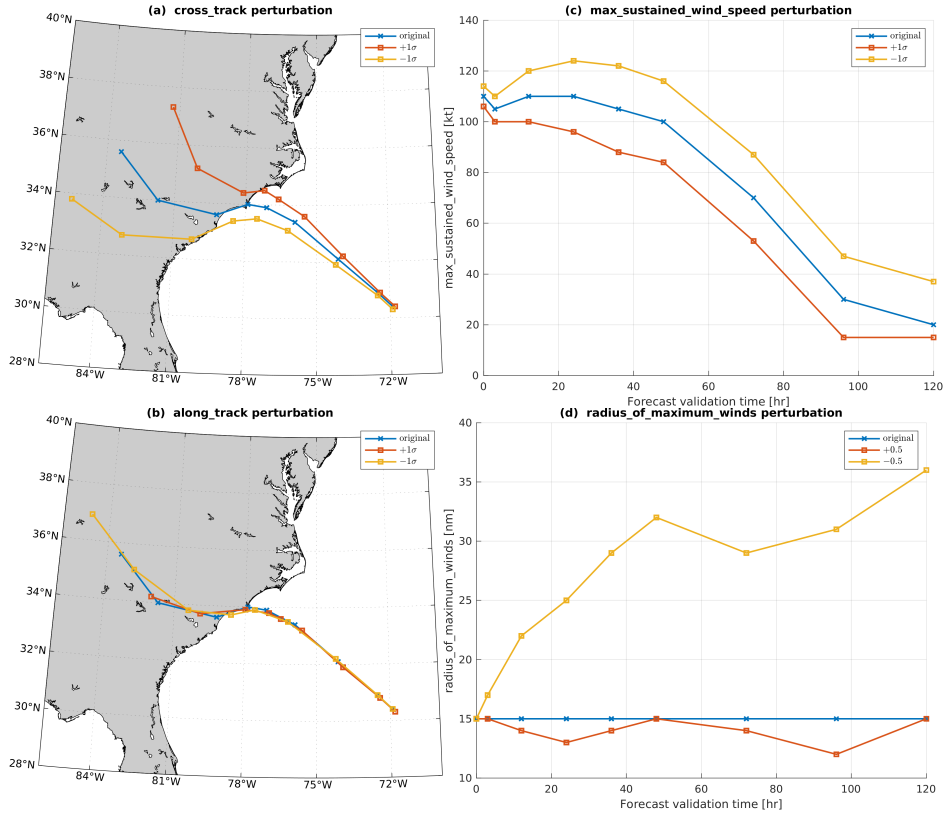
124 The hydrodynamic storm tide model is based on the ADvanced CIRCulation (ADCIRC; Luettich  
 125 and Westerink 2004) code, which solves the shallow water equations using the continuous Galerkin  
 126 finite-element method over an unstructured triangular mesh. ADCIRC is also the modeling engine  
 127 for the Global Surge and Tide Operational Forecast System ([https://registry.opendata.  
 128 aws/noaa-gestofs/](https://registry.opendata.aws/noaa-gestofs/)). In this study we use version 55 of ADCIRC (Pringle et al. 2021) including  
 129 both astronomical tides and atmospheric-driven surge, but without coupling to a wind-wave model  
 130 to capture the wave setup effect. Parametric representations of the TC vortex (based on track  
 131 advisories provided by the NHC) are used to construct surface wind and pressure forcing driving  
 132 storm surge in the ADCIRC model. Here, we use the classical symmetrical Holland vortex model

133 (Holland 1980) that is built directly into the ADCIRC code. The unstructured mesh used in this  
134 study encompasses the western North Atlantic and Gulf Coast region with 1.81 million vertices,  
135 and resolution ranges from roughly 200 m at the coast and overland up to a maximum of 46 km in  
136 the open ocean (Technology Riverside Inc. and AECOM 2015). The vertical datum is mean sea  
137 level (MSL) and the floodplain extends up to an elevation of 10 m above MSL. Manning's n friction  
138 coefficients, surface canopy coefficients, and surface directional effective roughness lengths based  
139 on land use data are used to account for surface roughness effects on the hydrodynamics and to  
140 modify the atmospheric forcing overland, respectively (Technology Riverside Inc. and AECOM  
141 2015).

#### 142 *b. Tropical Cyclone Perturbation*

143 In this study, the forecasted TC is perturbed according to historical NHC forecast error statistics  
144 of position, intensity, and size (Taylor and Glahn 2008). Positional members are perturbed based on  
145 estimated errors for cross-track (CT) and along-track (AT), whereby CT refers to a perpendicular  
146 offset of the forecast track and AT refers to a slowdown or speedup of the TC along the forecasted  
147 track. Intensity is described by both maximum sustained wind speed ( $V_{max}$ ) and minimum central  
148 pressure ( $P_c$ ), which is related through the  $B$  parameter of the Holland model (Holland 1980),  
149  $V_{max}^2 = B(P_b - P_c)/(e\rho)$ , where  $e$  is the base of natural logarithms,  $\rho$  is the air density, and  $P_b$   
150 is the background air pressure. We choose  $V_{max}$  as the independent perturbed variable and  $P_c$   
151 becomes a dependent variable based on keeping the  $B$  parameter consistent with the original  
152 forecast. The size of the TC is changed by perturbing the radius to maximum winds ( $R_{max}$ ).

156 Perturbation of CT, AT, and  $V_{max}$  across the whole forecast is achieved through a look-up table  
157 of historical NHC mean absolute forecast errors for certain lead times, distributed with the P-Surge  
158 model (Penny and Cangialosi 2019) and included here in Tables A1-A3. TCs are divided into  
159 three intensity bins based on the initial 0-hr  $V_{max}$  due to different error statistics between these  
160 categories. For instance, the mean absolute CT error is 11.6 nm at the 12-hr lead time and 27.8  
161 nm at the 48-hr lead time for the medium strength TC (50-95 kt). In this way if we perturb the  
162 CT variable by one 'mean absolute error' the track will be offset a perpendicular distance from  
163 the original position of 11.6 nm at the 12-hr mark and 27.8 nm at the 48-hr mark. The same idea  
164 is true for AT and  $V_{max}$ . The CT, AT and  $V_{max}$  errors are treated as Gaussian random variables



153 FIG. 2. Example perturbations of a NHC hurricane advisory (Hurricane Florence) along the forecast.  $\pm 1\sigma$   
 154 perturbations for the Gaussian distributed (a) CT, (b) AT, and (c)  $V_{max}$  errors, and  $\pm 0.5$  value perturbation of the  
 155 uniformly distributed (d)  $R_{max}$  errors.

165 ( $\mu = 0, \sigma = 1$ ) whereby the mean absolute error is defined as  $0.7979\sigma$  (Gonzalez and Taylor 2018).  
 166 A negative value will perturb the CT and AT in one direction and positive value in the other, and  
 167 similarly for  $V_{max}$  a negative value corresponds to an intensity underestimate of the forecast and  
 168 vice-versa for a positive value (see Fig. 2 for an illustration of  $\pm 1\sigma$  perturbations). Note that we  
 169 enforce [15,175] kt bounds on  $V_{max}$ .

170 For  $R_{max}$  the idea is similar to the other variables except that at each lead time the perturbation  
 171 is only from the initialized 0-hr value as NHC does not provide estimates of  $R_{max}$  along the  
 172 forecast. Unlike the other variables, the  $R_{max}$  error is bounded and treated as a random variable  
 173 with a uniform distribution ( $\in[-1,1]$ ) where the upper and lower error bounds at each lead time are  
 174 found through a look-up table with the TCs divided into five size bins based on the initialized 0-hr  
 175  $R_{max}$  (Table A4). We determined these upper and lower error bounds by linearly extrapolating

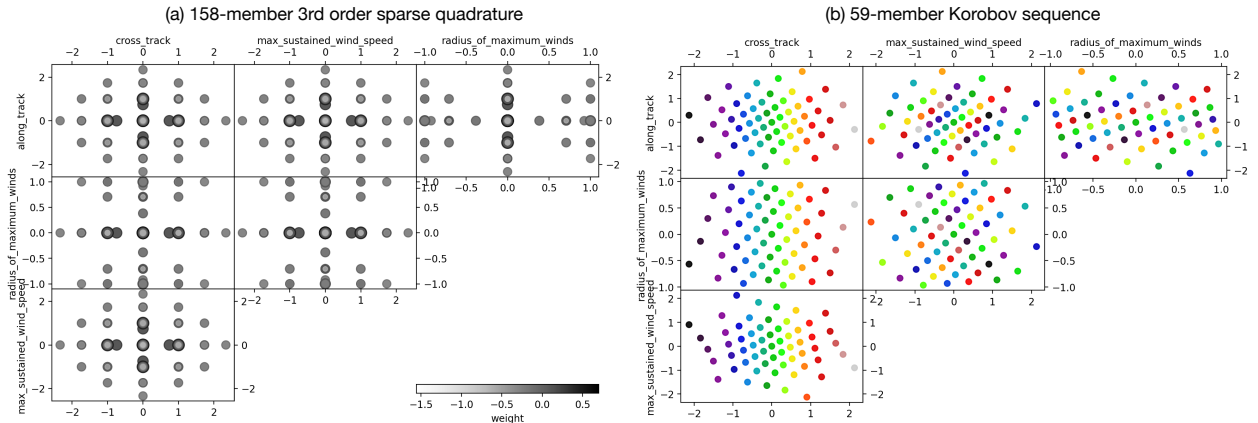


176 the values used by P-Surge at the 15th, 50th and 85th percentile to the 0th and 100th percentile.  
177 Referring to Table A4, an initially small storm is skewed towards having negative  $R_{max}$  forecast  
178 errors (becoming larger along the forecast), and vice-versa for initially large storms (see Fig. 2 for  
179 an illustration of  $\pm 0.5$  perturbations). Note that we enforce [5,200] nm bounds on  $R_{max}$ .

### 180 *c. Ensemble Generation*

181 An ensemble of TC forecasts is generated by sampling the random variables (CT, AT,  $V_{max}$  and  
182  $R_{max}$  errors) based on the probabilistic property of each variable equally, for forward uncertainty  
183 propagation analysis (section d). The idea is to build a surrogate model based on relatively few  
184 samples [ $O(10)$ ], which can be then trivially queried to generate the probabilistic forecast, as well  
185 as conduct a global sensitivity analysis or obtain a forecast given user-defined values of the TC error  
186 variables. In contrast, P-Surge employs an ad-hoc full factorial sampling methodology whereby 7  
187 perturbations of AT (slow to fast) and 3 of  $V_{max}$  and  $R_{max}$  (weak/large, “medium”, and strong/small)  
188 are used, along with enough perturbations of CT to cover 90% of the Gaussian distribution with  
189 spacing  $R_{max}$  at the 48-hr forecast (Gonzalez and Taylor 2018). Each possible permutation is used  
190 where each TC event is assigned a weight based on the combined probability and the probabilistic  
191 result is determined through summation of the weighted model results. This leads to 63 TC events  
192 per CT perturbation, or  $\sim 400$ -900 TC events based on 7 to 15 CT perturbations (Kyprioti et al.  
193 2021a), which would be prohibitive for HSOFS in a resource and time-limited environment.

198 We sample the variables using a quasi-random low-discrepancy sequence, of which several  
199 are available in the chaospy python package (Feinberg and Langtangen 2015) employed by  
200 `EnsemblePerturbation`, including widely-used Sobol and Halton types. Here, we recommend  
201 the use of the Korobov sequence (Korobov 1959) because the random variables are sampled  
202 symmetrically about zero and cover a predictable range across all variables for any given sample  
203 size, which is not the case for the other chaospy sequence implementations. The benefit of  
204 such low-discrepancy sequences is avoidance of the “curse of dimensionality” that is associated  
205 with quadrature integration, which the P-Surge methodology could be viewed as a subset of. For  
206 instance, 3rd order quadrature integration for the four-dimensional problem requires  $4^4 = 256$   
207 samples, as all possible permutations of just four perturbations of each variable is used. Smolyak  
208 sparse grid quadrature can be used to alleviate the issue, although  $\sim 150$  samples are still required

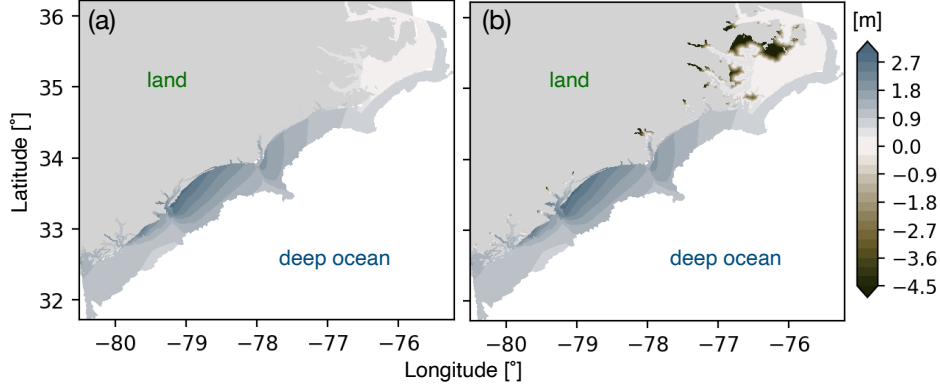


194 FIG. 3. Perturbed values of  $\lambda$  across the four dimensional space for (a) sparse quadrature and (b) Korobov  
 195 sequence sampling methodologies. Sparse quadrature has variable weights attached to each perturbation sample  
 196 (indicated by marker gray-scale and size) while the Korobov sequence is equally weighted (marker colors are  
 197 used to distinguish unique samples across the panels).

209 for 3rd order quadrature. Instead, a low-discrepancy sequence can be used to sample each variable  
 210 more densely without imposing the condition of using all permutations across the four-dimensional  
 211 space, as illustrated in Fig. 3. Later we show that these low-discrepancy sequences with sample  
 212 size  $O(10)$  can be used to generate a surrogate model that, due to enhanced regression techniques,  
 213 are indeed of improved quality over the sparse quadrature.

#### 214 d. Forward Uncertainty Propagation

215 We wish to know a probabilistic form of the model output  $Z = f(\lambda, \mathbf{x})$  (maximum water surface  
 216 elevation in  $\mathbf{x}$ ) which is dependent on the set of input TC error parameters,  $\lambda = (\text{CT}, \text{AT}, V_{max},$   
 217  $R_{max})$ . However, the underlying hydrodynamic storm tide model is too computationally expensive  
 218 to sample a large number of times to properly understand the uncertainty and sensitivity of  $Z$   
 219 to the TC error variables. Therefore, we employ a surrogate approximation  $g(\lambda, \mathbf{x}) \approx f(\lambda, \mathbf{x})$   
 220 through PC theory (Sargsyan 2017; Sochala et al. 2020), which is constructed from a training set  
 221 (section c). The resulting PC surrogate model is then a parametric representation of  $Z$  which can  
 222 be trivially sampled a large number of times, and from which moments and global sensitivities can  
 223 be analytically extracted. To ensure a suitable training set for generating an accurate PC surrogate



232 FIG. 4. Example of maximum water surface elevation extrapolation over dry mesh points for a Hurricane  
 233 Florence training set member. (a) water surface elevations before extrapolation, (b) water surface elevations after  
 234 extrapolation. Note that mesh points with ocean depths greater than 25 m have been removed from the dataset.

224 (which requires a degree of smoothness) we manipulate  $Z$  and apply a dimensionality reduction  
 225 technique for computational efficiency. These methodologies are outlined in the rest of this section.

## 226 1) MODEL OUTPUT MANIPULATION

227 There are two related difficulties that we encounter here with using the maximum water surface  
 228 elevation for training the PC surrogate: 1) Some mesh points are inundated during some TC events  
 229 and not in others, and; 2) water surface elevation cannot physically go below the ground elevation  
 230 (water depth must be positive or zero) but the surrogate model cannot be easily constrained to  
 231 prevent a physically unrealistic negative depth prediction.

235 In the first problem the intuitive solution is to set  $Z$  to that of the ground elevation for mesh  
 236 points that are not inundated (denoted ‘NaN’) in a certain TC event. However, this does not  
 237 distinguish between TC events where the mesh point of concern might be have been close to being  
 238 inundated or very far from being inundated, resulting in a poor PC fit for predicting inundation.  
 239 The solution we propose is to artificially extrapolate  $Z$  from wet mesh points over dry mesh points  
 240 for the PC surrogate training purposes (Fig. 4). Here, we use an inverse-distance weighting (IDW)  
 241 extrapolation (Plumlee et al. 2021) with a hydraulic head loss factor,

$$\tilde{Z}_d = \frac{\sum_{w=1}^k (Z_{dw} - h_{dw}) D_{dw}^{-p}}{\sum_{w=1}^k D_{dw}^{-p}} \quad (1)$$

242 where  $\tilde{Z}_d$  is the artificial maximum water surface elevation of the dry mesh point  $d$ ,  $Z_{dw}$  is the  
 243 maximum water surface elevation of the  $w^{th}$  closest wet mesh point to  $d$ ,  $D_{dw}$  is the distance from  
 244 mesh point  $w$  to  $d$ ,  $p$  is the IDW extrapolation order,  $k$  is the number of nearest neighbors to use  
 245 for extrapolation, and  $h_{dw}$  is the head loss (Rucker et al. 2021) from mesh point  $w$  to  $d$ ,

$$h_{dw} = D_{dw} f_f \quad (2)$$

246 where  $f_f$  is a hydraulic friction factor. This methodology is similar to the weighted  $k$  nearest  
 247 neighbour pseudo-surge methodology used by Kyprioti et al. (2021c), in which there are four free  
 248 parameters of the weighting scheme that require calibration. Ostensibly, the head loss factor we use  
 249 here adds some physical meaning to the extrapolation. The factor  $f_f$  can be related to Manning's  
 250 equation like in Rucker et al. (2021), requiring the Manning's  $n$  friction coefficient, flow velocity,  
 251 and flow depth. These flow quantities are not available once  $Z$  is extrapolated over the dry regions  
 252 so we simply view this relation in terms of guiding  $f_f$  to a physically reasonable value. In section 2  
 253 we compare values for  $f_f$ ,  $k$ , and  $p$  and how they affect surrogate model prediction accuracy.

254 A possible solution to the second problem is to build the surrogate model based on  $\log(H)$   
 255 (Plumlee et al. 2021), where  $H$  is the simulated maximum water depth, which is physically always  
 256 positive, guaranteeing that the surrogate prediction will be positive. However, an issue we find  
 257 here is that our water surface elevation extrapolation technique proposed above leads to artificial  
 258 negative water depths for otherwise dry mesh points in the training set. Therefore,  $H$  in the training  
 259 data would need to be modified to be positive for these points by adding a constant, which can be  
 260 subtracted back from the surrogate prediction. Of course, this means that the surrogate model can  
 261 actually predict a negative real water depth for such dry mesh points, just as was provided to it  
 262 for training. We test the accuracy of constructing the surrogate in log-space versus linear-space in  
 263 section 2.

## 264 2) DIMENSIONALITY REDUCTION AND POLYNOMIAL CHAOS SURROGATE

265 Building a surrogate model for all HSOFS mesh points (1.81 million), or even a subset of points  
 266 around hurricane landfall [ $O(10^5)$ ], would be prohibitively expensive, therefore we seek a method  
 267 to reduce the dimensionality of the problem. Such dimensionality reduction is common practice  
 268 and has been used for building other surge surrogate models (e.g., Jia et al. 2016; Sochala et al.

269 2020; Kyprioti et al. 2021b; Lee et al. 2021). Here, we achieve dimensionality reduction via  
 270 Karhunen-Loève (KL) expansions which are then coupled with PC surrogates.

271 As before, our model output of maximum water surface elevations,  $Z = f(\boldsymbol{\lambda}, \boldsymbol{x})$  is dependent on  
 272 the set of input TC error parameters,  $\boldsymbol{\lambda}$  and is spatially varying with  $\boldsymbol{x}$ . The KL expansion can be  
 273 written as,

$$Z = f(\boldsymbol{\lambda}, \boldsymbol{x}) = \bar{f}(\boldsymbol{x}) + \sum_{j=1}^L \xi_j(\boldsymbol{\lambda}) \sqrt{\mu_j} \phi_j(\boldsymbol{x}) \quad (3)$$

274 in terms of uncorrelated, zero-mean, unit-variance random variables  $\xi_j(\boldsymbol{\lambda})$  and eigenvalue-  
 275 eigenfunction pairs  $(\mu_j, \phi_j(\boldsymbol{x}))$  of the covariance,

$$C(\boldsymbol{x}, \boldsymbol{x}') = E_{\boldsymbol{\lambda}}[(f(\boldsymbol{\lambda}, \boldsymbol{x}) - \bar{f}(\boldsymbol{x}))(f(\boldsymbol{\lambda}, \boldsymbol{x}') - \bar{f}(\boldsymbol{x}')))] \quad (4)$$

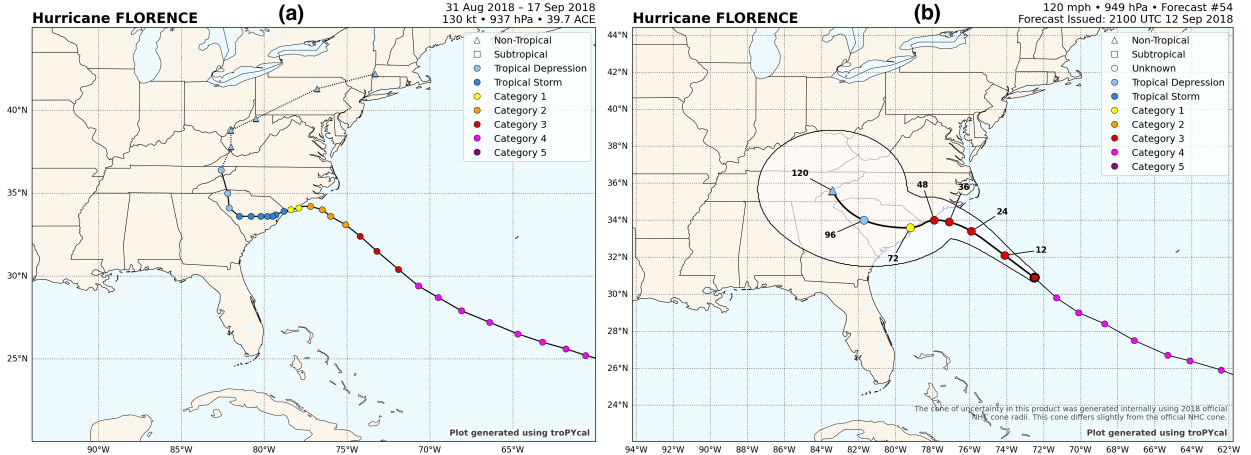
276 truncated at eigenvalue  $L$  ( $\ll$  dimensions of  $\boldsymbol{x}$ ) that explains a user-defined level of variance. The  
 277 expectation  $E_{\boldsymbol{\lambda}}$  indicates averaging across parameter  $\boldsymbol{\lambda}$ , as does the bar symbol, i.e.,  $\bar{f}(\boldsymbol{x}) =$   
 278  $E_{\boldsymbol{\lambda}}[f(\boldsymbol{\lambda}, \boldsymbol{x})]$ . The forward uncertainty propagation problem therefore reduces to seeking a function  
 279 approximation for the KL coefficient functions  $\xi_j(\boldsymbol{\lambda})$ , for which we employ a PC form here,

$$\xi_j(\boldsymbol{\lambda}) \approx \sum_{k=0}^K c_{jk} \Psi_k(\boldsymbol{\xi}) \quad (5)$$

280 where  $\Psi_k(\boldsymbol{\xi})$  are multivariate orthogonal polynomials with respect to the PDF of the stochastic  
 281 germ  $\boldsymbol{\xi}$ , which is a vector with elements being standard random variables that are chosen according  
 282 to the expected PDF of the corresponding element of  $\boldsymbol{\lambda}$ , i.e., Gauss-Hermite for CT, AT and  $V_{max}$   
 283 errors and the Legendre-Uniform for  $R_{max}$  errors in this study. Finally, by substituting the PC  
 284 equations (5) into the KL expansion (3) and switching the summations we arrive at the following  
 285 joint KL-PC surrogate expansion,

$$Z = f(\boldsymbol{\lambda}, \boldsymbol{x}) \approx g(\boldsymbol{\lambda}, \boldsymbol{x}) = \sum_{k=0}^K c_k(\boldsymbol{x}) \Psi_k(\boldsymbol{\xi}), \quad \text{where} \quad (6)$$

$$c_k(\boldsymbol{x}) = \delta_{k0} \bar{f}(\boldsymbol{x}) + \sum_{j=1}^L c_{jk} \sqrt{\mu_j} \phi_j(\boldsymbol{x}) \quad (7)$$



303 FIG. 5. Hurricane Florence 2018 track and intensity information; (a) best-track hindcast, (b) NHC forecast  
 304 advisory 48-hr prior to landfall.

286 and  $\delta$  is the Kronecker delta. Given the PC coefficients  $c_k(\mathbf{x})$ , moments and global sensitivity  
 287 indices can be analytically extracted from the KL-PC surrogate expansion (Sargsyan 2017; Sargsyan  
 288 et al. 2021).

289 In our python implementation of the joint KL-PC surrogate expansion in  
 290 EnsemblePerturbation, the scikit-learn Principal Component Analysis (PCA) class  
 291 is used to perform the KL decomposition, and the chaospy (Feinberg and Langtangen 2015) pack-  
 292 age is used to perform the PC expansion, utilizing linear regression models from scikit-learn.  
 293 These regression models are introduced and assessed in section 3.

294 *e. Experimental Design*

295 HSOFS is primarily concerned with forecasts of hurricane surge and inundation near (<48-  
 296 hr) landfall. To provide a rigorous test of the methodologies (section a), in this study we explore  
 297 results for the 48-hr lead time NHC advisory for three historical US hurricanes; Irma 2017, Florence  
 298 2018, and Laura 2020 (e.g., Fig. 5b). The storm tide model is spun-up from a quiescent state with  
 299 astronomical tides and best-track hindcast forcing (e.g., Fig 5a) for 7-days prior to the forecast. For  
 300 each hurricane we analyze a subset of the model based on mesh points that have ocean depths  $\leq 25$   
 301 m and that fall within the 34-kt wind speed swath of the NHC advisory. The spatial dimension of  
 302 this subset is reduced through KL decomposition with truncation at the 99.99% variance level.

305 In the first set of methodology experiments (section a1) we compare regression models, sampling  
306 methods, and sample size for construction of the surrogate model. In these experiments we conduct  
307 the analysis in linear-space only on mesh points that are inundated across all TC events to avoid  
308 complicating the experiment with the water surface elevation extrapolation over dry mesh points.  
309 First, we compare sparse quadrature integration (158 samples) to a 59-member Korobov sequence  
310 using different `scikit-learn` regression models for constructing the surrogate model. Then using  
311 the best regression model from that experiment we compare Korobov sequences with sample sizes  
312 of 19, 39, and 59 which cover 90.0%, 95.0% and 96.7% of the distributions of  $\lambda$ , respectively.  
313 In the second set of methodology experiments (section a2) we conduct the analysis on all mesh  
314 points in the subset, in both log-space versus linear-space, and with varying parameters of the  
315 water surface elevation extrapolation method. Note that in all cases we use 3rd order PCs which  
316 we found to be the only reliable PC order; 2nd order PCs are not flexible enough while the desired  
317 size of the training set is too small to allow for higher-order four-dimensional PCs to be constructed  
318 accurately.

319 To validate the surrogate model we use a 128-member validation set for each storm where  $\lambda$  is  
320 randomly chosen from their distributions for each validation member. To evaluate the accuracy  
321 of the surrogate model we compute the root-mean-square-error (RMSE) across the mesh points  
322 between the model simulation and the surrogate model for a single validation member. To compare  
323 the surrogate model across all validation members we plot cumulative distribution functions (CDFs)  
324 of the RMSE and compute the two-sided t-test statistic of the RMSE distribution between two  
325 surrogate models. In the second set of experiments we also compute the percentage of mesh points  
326 that are falsely classified as wet or dry in the surrogate model prediction.

327 Finally, using the recommend methodology based on the experiments we construct joint KL-  
328 PC surrogate models for each hurricane to produce example products of an ensemble HSOFS,  
329 i.e., global sensitivities of  $Z$  with respect to  $\lambda$ , and exceedance probability maps (section b).  
330 The probabilistic predictions are compared to simulated model results of the best-track hindcast  
331 hurricane forcing. The reliability of the probabilistic prediction is assessed by comparing the  
332 fraction of elevation exceedances in the best-track results above the height of the given exceedance  
333 probability. Here, mesh points where the predicted exceedance elevation is NaN (dry) are ignored  
334 in the computation, while a NaN in the best-track simulation is set to ground elevation. In addition,

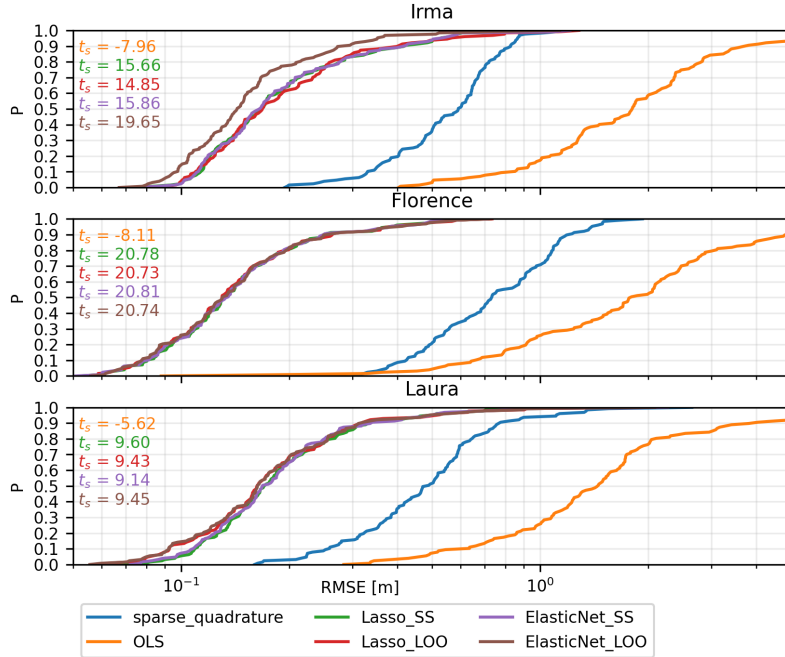


FIG. 6. CDF curves of the surrogate model RMSE across all validation members (128 samples) for the three hurricane forecasts (Irma, Florence, Laura), using a 158-member sparse quadrature training set and a 59-member Korobov sequence training set with different regression methodologies.  $t_s$  is the two-sided t-test statistic between the sparse quadrature RMSE and the Korobov RMSEs corresponding to the colors in the legend (largest positive value indicates smallest average RMSE).

inundation area is compared, where we expect the median (50% exceedance probability) prediction to be similar to that of the best-track.

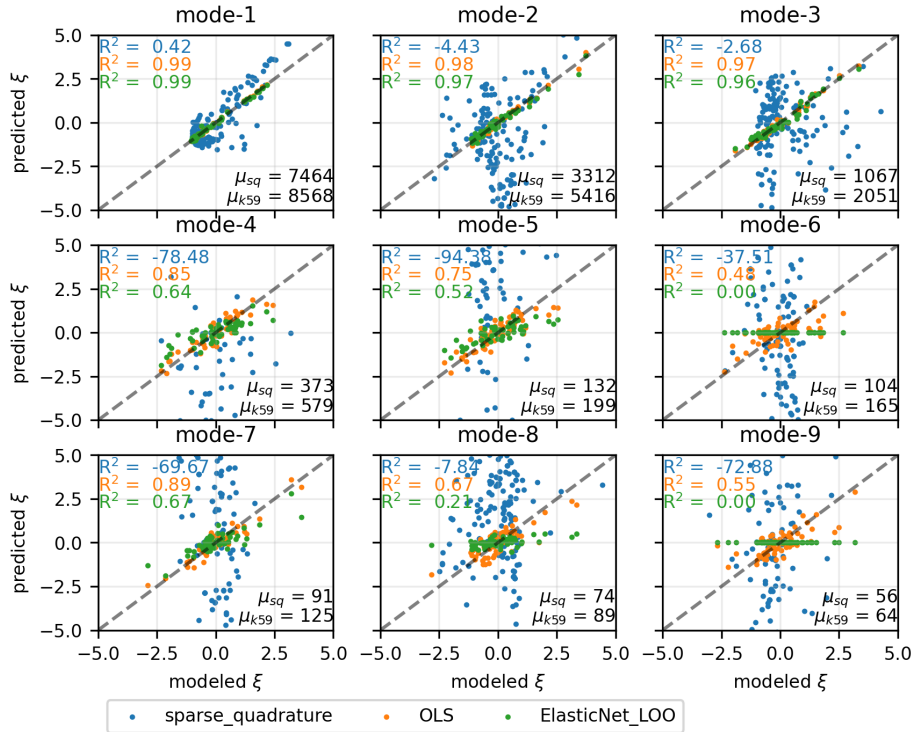
### 3. Results

#### a. Methodology Experiments

##### 1) REGRESSION

Results across all three storms show that fitting the surrogate model using (sparse) quadrature integration is superior to Ordinary Least Squares (OLS) linear regression but far inferior to penalized linear regression from the 59-member Korobov sequence (Fig. 6). For this test we compare Lasso and Elastic Net regression that uses  $\ell_1$ -norm regularization and combined  $\ell_1$ -norm and  $\ell_2$ -norm regularization with equal weighting, respectively.  $\ell_1$ -norm regularization penalizes non-zero coef-





360 FIG. 7. Comparison of the predicted (surrogate) and modeled KL training parameters,  $\xi$ , of the top nine KL  
 361 modes for the Hurricane Florence forecast.  $R^2$  is the coefficient of determination of the surrogate prediction  
 362 corresponding to the colors in the legend.  $\mu$  indicates the eigenvalues of each KL mode evaluated from the sparse  
 363 quadrature (sq) and 59-member Korobov sequence (k59) training data.

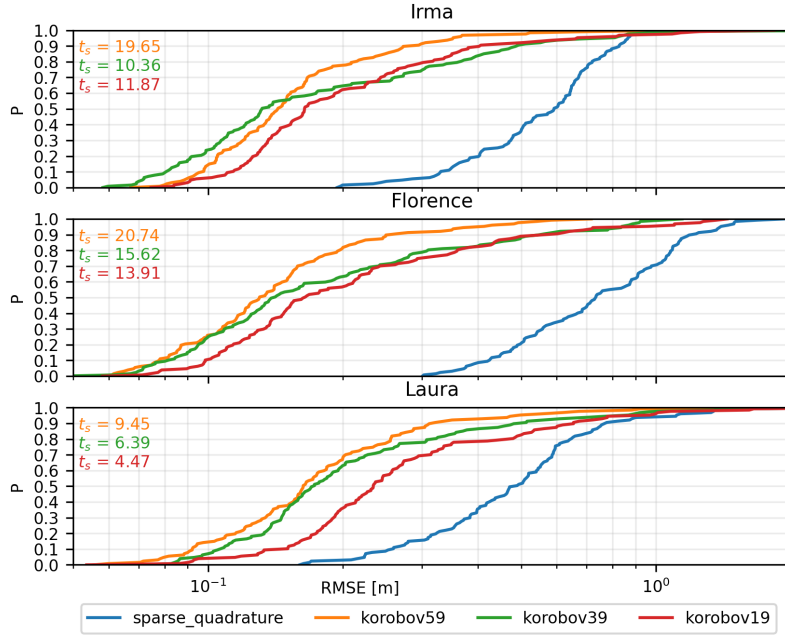
350 ficients to form sparse models, and  $\ell_2$ -norm regularization penalizes the size of the coefficients to  
 351 form smooth models. These linear regression models avoid overfitting the KL parameters,  $\xi$ , in the  
 352 training set which we see for OLS regression (Fig. 7), resulting in poor validation. As a rule, the  
 353 penalized regression gives a poorer fit to the KL training parameters than for the OLS regression,  
 354 although correlation is generally higher for lower modes. On the other hand, sparse quadrature  
 355 gives a very poor fit to the KL training parameters, which indeed leads to poor validation relative  
 356 to the penalized regression. Sparse quadrature projection is known to perform poorly for noisy  
 357 data due to the presence of negative quadrature weights and amplification of small errors during  
 358 PC construction (Sargsyan 2012). Nevertheless, sparse quadrature does validate better than OLS  
 359 regression.

364 With both Lasso and Elastic Net we use built-in cross-validation estimators to automatically  
365 select the best penalization weight and return the most robust fit to the data. Shuffle-Split (SS) and  
366 Leave-One-Out (LOO) cross-validation estimators are used to automatically divide up the overall  
367 59-member training set into training and validation subsets during the regression fitting process.  
368 Results show that overall there are relatively small differences between the four combinations of  
369 cross-validators and regularization strategies, although ElasticNet\_LOO notably outperforms for  
370 Irma, as well as having a strong performance for both Florence and Laura. Therefore, we decide to  
371 use ElasticNet\_LOO for the remainder of this paper, remarking also that LOO is attractive because  
372 there are no parameter choices to be made, while SS requires choosing the relative size of the  
373 training and validation subsets (we used the `scikit-learn` default options here), in which the  
374 optimal choice may differ with sample sizes and storms. Furthermore, for Elastic Net regression  
375 we can also use cross-validation to select the optimal weighting between  $\ell_1$  and  $\ell_2$  penalties if so  
376 desired, although we used 0.5 (equal weighting) here for simplicity.

382 When reducing the training sample size using the Korobov sequence, validation performance  
383 remains notably superior to the sparse quadrature baseline, but does degrade as expected (Fig. 8).  
384 The 39-member sequence performs about as well as, or better in the case of Irma, than the 59-  
385 member sequence in the lower half distribution but noticeably worse in the upper half distribution.  
386 While the 19-member sequence performs similarly to the 39-member sequence in the upper half  
387 distribution. If for practical purposes, we select the 90th percentile of the RMSE as an arbitrary  
388 measure of performance (RMSE<sub>90</sub>), for all three storms the 59-member sequence has an RMSE<sub>90</sub>  
389 accuracy of approximately 0.3 m. Whereas, the RMSE<sub>90</sub> accuracy is approximately 0.5 m for  
390 the 39-member sequence and 0.5-0.7 m for the 19-member sequence. Therefore, in section 2 we  
391 choose the 59-member Korobov sequence as it provides about twice the accuracy of the smaller  
392 sample sizes under this assessment.

## 393 2) WATER SURFACE ELEVATION EXTRAPOLATION AND LOGARITHMIC TRANSFORMATION

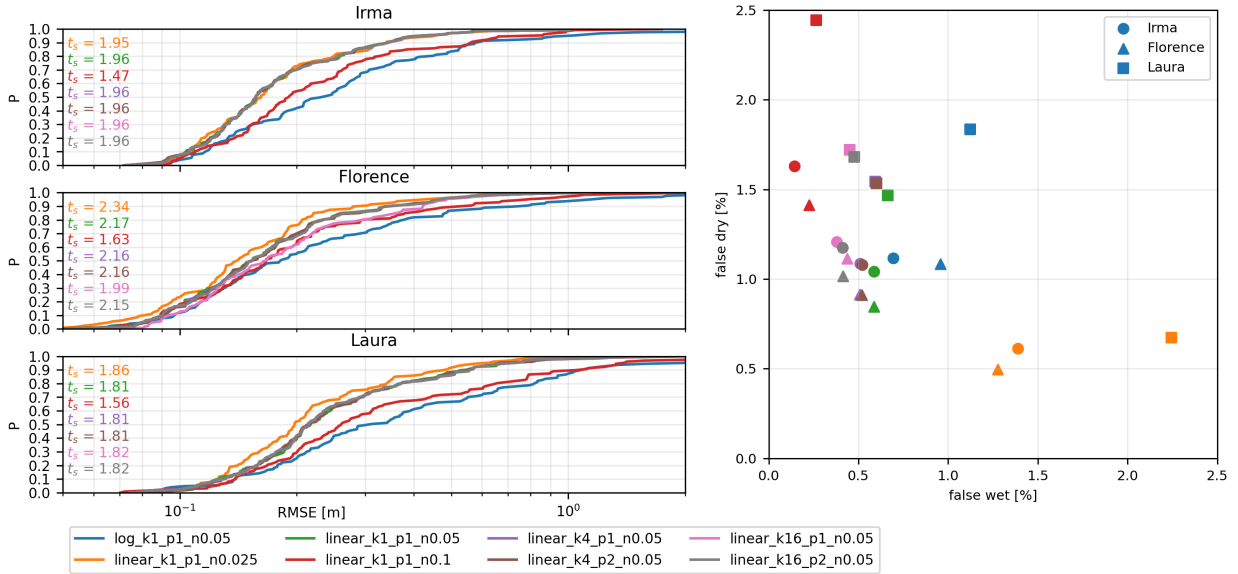
394 Here we vary the number of IDW neighbors ( $k = [1, 4, 16]$ ), IDW order ( $p = [1, 2]$ ), and  
395 the friction factor ( $f_f = [0.0001, 0.0004, 0.0016]$ ) for the water surface elevation extrapolation  
396 over dry mesh points. The friction factor values are derived through the Manning's relation,  
397  $f_f = n^2 U_f^2 / H_f^{4/3}$ , with  $n = [0.025, 0.05, 0.1]$   $\text{sm}^{-1/3}$  (Manning's  $n$  coefficient),  $H_f = 1$  m (flow



377 FIG. 8. CDF curves of the surrogate model RMSE across all validation members (128 samples) for the three  
 378 hurricane forecasts (Irma, Florence, Laura), using 19-, 39-, and 59-member Korobov sequence training sets  
 379 with ElasticNet\_LOO regression.  $t_s$  is the two-sided t-test statistic between the sparse quadrature RMSE and  
 380 the Korobov RMSEs corresponding to the colors in the legend (largest positive value indicates smallest average  
 381 RMSE).

398 depth) and  $U_f = 0.4 \text{ ms}^{-1}$  (flow velocity). The surrogate model for the  $k = 1, p = 1, n = 0.05$  case  
 399 is computed in log-space as well as linear-space. Results show that surrogate model accuracy is  
 400 worse in terms of both RMSE and false wet/dry classification than in linear-space (Fig. 9). We  
 401 also tried surrogate model generation in log-space only for mesh points that are always wet in  
 402 the training set (like in section a1), which did indeed provide an improvement to the false dry  
 403 classification percentage compared to linear-space (not shown). However, it would appear that  
 404 when extrapolation is used over dry points and negative depths are introduced to the training set,  
 405 this benefit disappears. Therefore, we only show the other extrapolation parameter experiment  
 406 results in linear-space.

415 Surrogate model RMSE tends to increase with the friction factor, especially for  $n = 0.1$  (Fig. 9).  
 416 Performance for  $n = 0.025$  and  $0.05$  are similar except for Laura, in which the smaller  $n = 0.025$  is  
 417 clearly superior. False wet/dry classifications follow a clear pattern where the surrogate model with



407 FIG. 9. Comparison of the surrogate model accuracy for different extrapolation parameters and construction  
 408 in linear-space versus log-space.  $k$  and  $p$  in the legend are parameters from the extrapolation Eq. (1).  $n$  is the  
 409 Manning’s  $n$  coefficient used to compute  $f_f$  in Eq. (2). Results are shown across all validation members (128  
 410 samples) for the three hurricane forecasts (Irma, Florence, Laura), using 59-member Korobov sequence training  
 411 sets with ElasticNet\_LOO regression. Left: RMSE CDF curves, where  $t_s$  is the two-sided t-test statistic between  
 412 the log-space RMSE and the linear-space RMSEs corresponding to the colors in the legend (largest positive value  
 413 indicates smallest average RMSE). Right: Total percentage of mesh points across all validation members with a  
 414 false wet/dry classification.

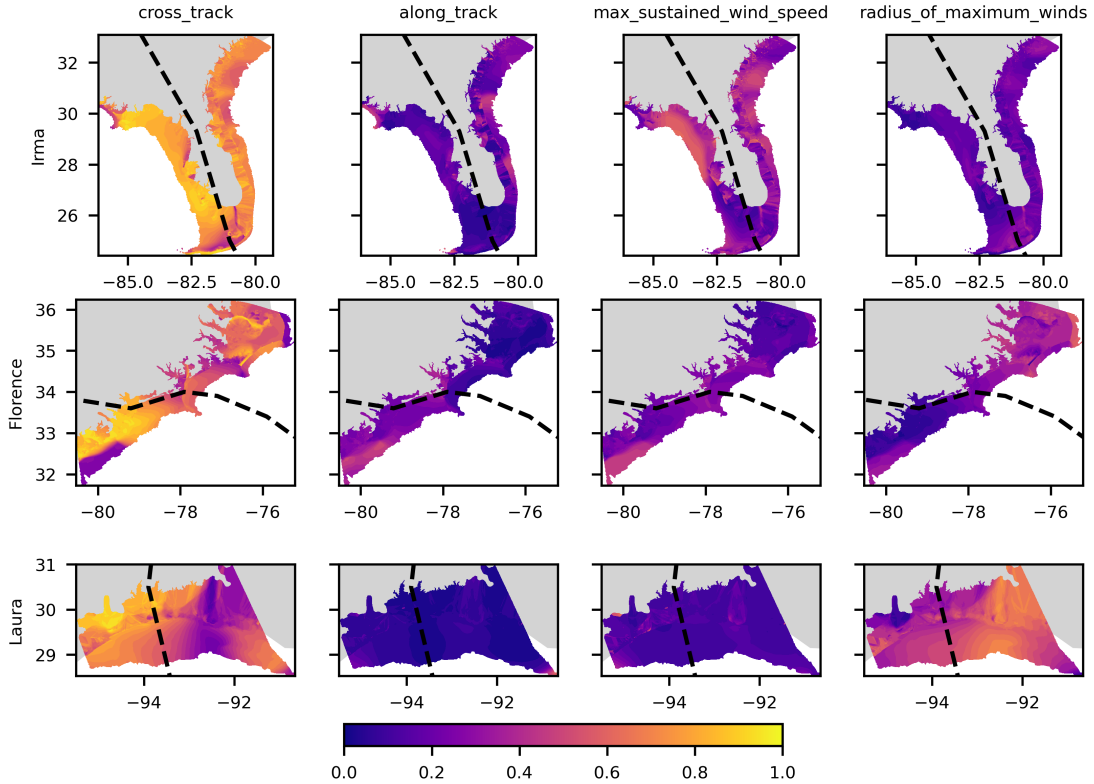
418 a larger friction factor gives more false dry predictions but fewer false wet predictions, and vice-  
 419 versa for a smaller friction factor. In addition to the fact that the smaller friction factor produces  
 420 smaller RMSEs, we prefer the smaller  $n = 0.025$ , since from an emergency management standpoint  
 421 it would generally be considered preferable to be biased towards false wet classifications. As for the  
 422 IDW extrapolation parameters, using more neighbors and going to second-order does not provide  
 423 any discernible benefit to nearest neighbor ( $k = p = 1$ ). As such, and for a preference towards  
 424 simplicity, nearest neighbor using  $n = 0.025$  is selected for presentation of the results in section b.

425 *b. Probabilistic Predictions and Global Sensitivities*

426 Maps of the sensitivities and probabilistic predictions extracted from the best surrogate model  
427 setup from section a (59-member Korobov sequence training set with ElasticNet\_LOO regression  
428 and extrapolation using  $n = 0.025$  ( $f_f = 0.0001$ ) and  $k = p = 1$ ), are shown here to demonstrate  
429 the product output. First, total effect sensitivity indices of  $\lambda$  plotted in Fig. 10 indicate that the  
430 CT error is the most sensitive variable across all storms and over most of the region. The CT  
431 sensitivity tends to be smaller on the right-hand side of the forecasted track, likely since right-hand  
432 side wind speeds are supported by the hurricane forward speed. The importance of the other  
433 error variables is somewhat storm and location dependent.  $V_{max}$  is the second-most important  
434 for Irma, while  $R_{max}$  is for Laura. Florence is approximately equally sensitive to AT,  $V_{max}$ , and  
435  $R_{max}$ . This information could be used in conjunction with weather forecaster assessments of  
436 variable uncertainty to determine which regions have higher storm surge and inundation prediction  
437 uncertainty for the particular storm.

442 Second, 10%, 50%, and 90% exceedance probabilities of the maximum water surface elevations  
443 are shown in Fig. 11, illustrating how the surrogate model can predict changes to both water  
444 levels and inundation extents across the distribution. Indeed, over most of the domain the 50%  
445 exceedance probability tends to show a closer match to the best-track hindcast than the 10% or  
446 90% probabilities. However, as expected, in the regions where a large or small maximum water  
447 surface elevation occurs in the hindcast, the match appears closer to the 10% and 90% exceedance,  
448 respectively.

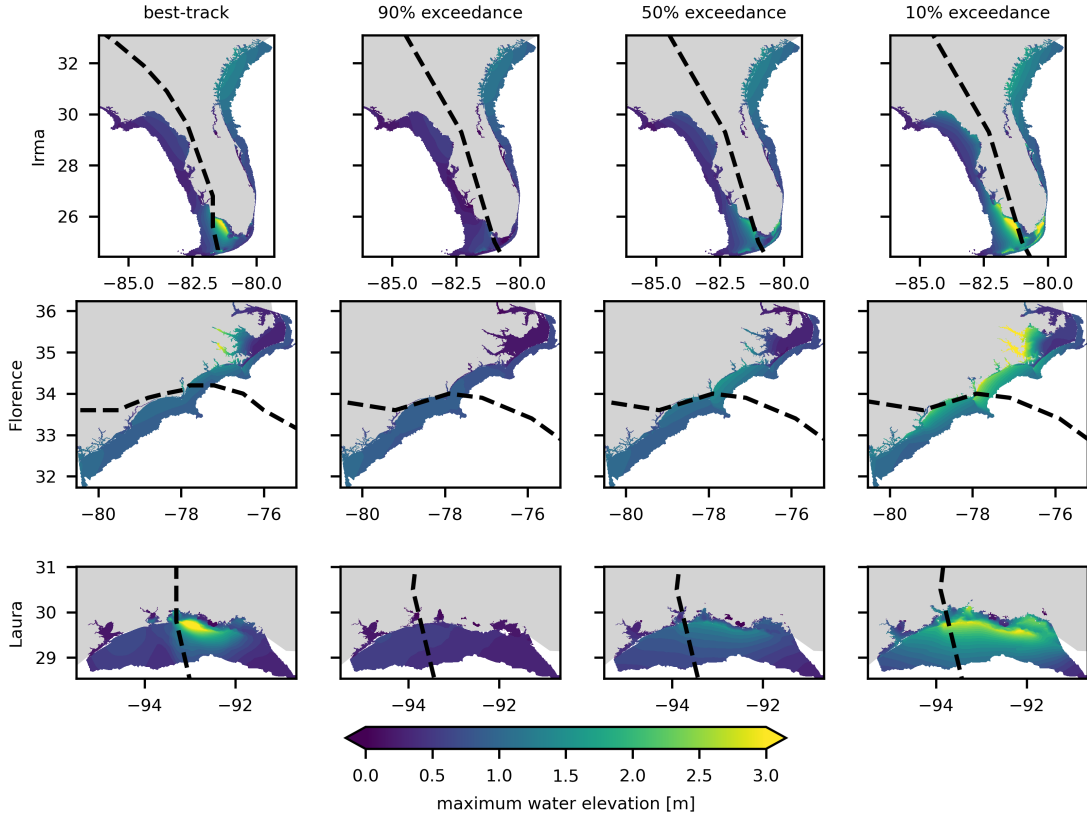
453 More quantifiably, reliability assessments show that surrogate model probabilistic forecast has  
454 generally reasonable accuracy for all three storms (Fig. 12). The 10% exceedance for all three storms  
455 is greater than that of the best-track indicating a high-bias at this extreme end of the distribution.  
456 In comparison, at the low end (towards 90% exceedance), the Irma and Florence predictions are  
457 biased low, while Laura is consistently biased high over the whole distribution. This may be  
458 related to the higher sensitivity to  $R_{max}$  for Laura than the other storms – best-track results show  
459 a relatively localized high water surface elevation region. Notably,  $R_{max}$  is treated differently than  
460 the other error variables and is difficult to measure and forecast, motivating alternative treatment  
461 for storm size in the future.



438 FIG. 10. Total effect sensitivity indices of maximum water surface elevation ( $Z$ ) to  $\lambda$  (CT, AT,  $V_{max}$ ,  $R_{max}$ )  
 439 for the three hurricane forecasts (Irma, Florence, Laura). The dashed black line is the track of the NHC forecast  
 440 advisory 48-hr prior to landfall. Note that mesh points with ocean depths greater than 25 m have been removed  
 441 from the dataset.

466 Interestingly however, in terms of inundation area, the 50% exceedance probability was close to  
 467 the best-track hindcast for Laura as was for Irma (Fig. 13). Though, the inundation area for Laura  
 468 is more sensitive overall to the choice of exceedance probability than the other storms, highlighting  
 469 the larger uncertainty for this hurricane. The inundation area for Florence is underestimated at  
 470 the 50% exceedance, only matching the best-track area at the 30% exceedance. Notably, the  
 471 10% exceedance inundation area for Florence is about a factor of 2 greater than the best-track,  
 472 demonstrating large uncertainty at the lower probability end of the distribution.

477 Comparing results for surrogate models trained on smaller sample sizes of the Korobov sequence  
 478 show remarkably consistent results in terms of reliability and inundation area across all storms  
 479 (Figs. 12,13). Generally, reliability of the surrogate model trained on the 19-member Korobov

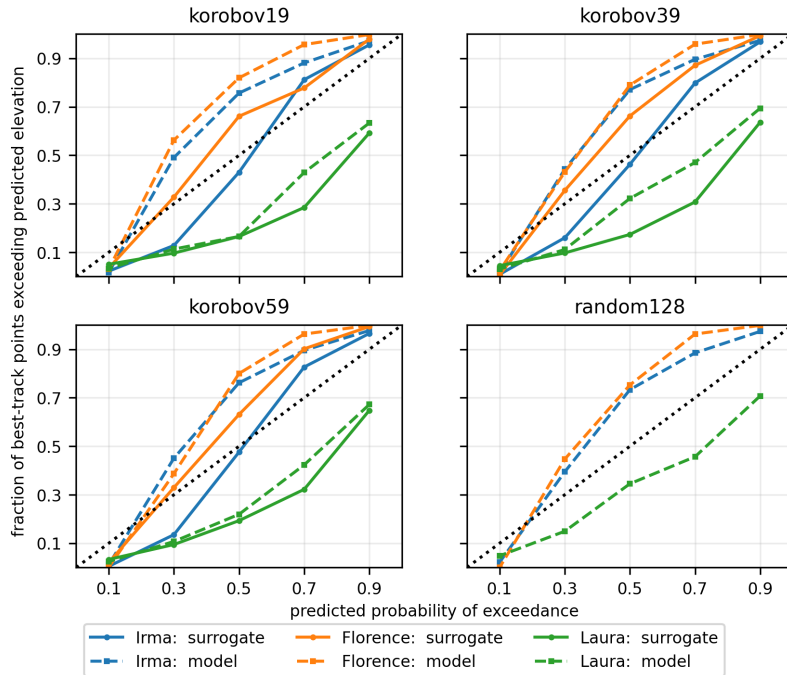


449 FIG. 11. Best-track hindcast and probabilistic predictions of maximum water surface elevations (10%, 50%,  
 450 and 90% exceedance probabilities) for the three hurricane forecasts (Irma, Florence, Laura). The dashed black  
 451 line is the best-track for the left hand side panels or the track of the NHC forecast advisory 48-hr prior to landfall  
 452 in the other panels. Note that ocean depths greater than 25 m have been removed from the dataset.

480 sequence is at least as good as for 59 members. For Laura the reliability increases slightly as the  
 481 sample sizes increase, and the best result is found from direct empirical evaluation of the 128-  
 482 member validation set. For Irma and Florence, the surrogate model clearly shows improvement  
 483 in reliability over direct empirical evaluation of the training and validation sets, highlighting its  
 484 potential added value.

#### 485 4. Discussion

486 The framework developed here has demonstrated that reliable probabilistic predictions of storm  
 487 tide elevations and inundation can be achieved by training a KL-PC surrogate model on just  $O(10)$   
 488 perturbed storm events using low-discrepancy Korobov sequences. The use of the surrogate model

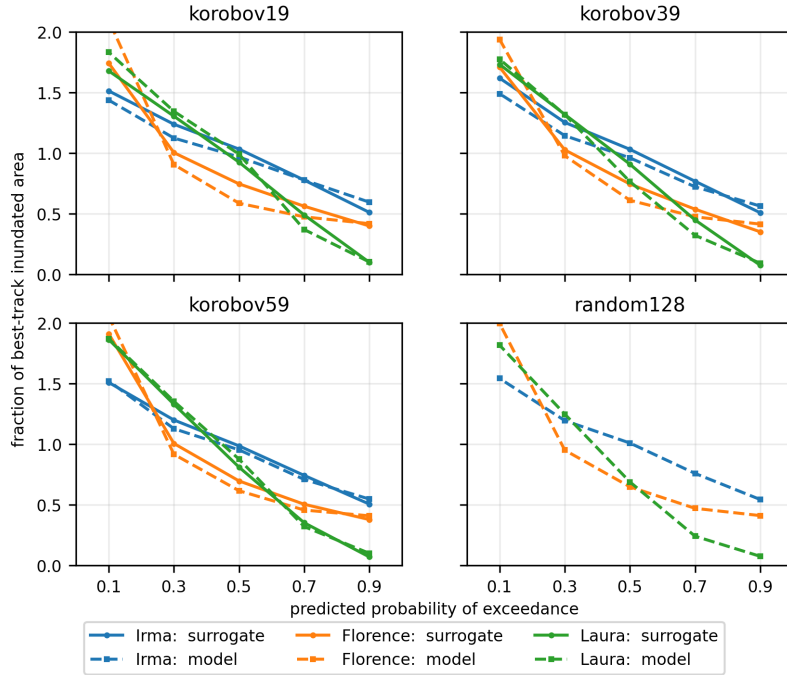


462 FIG. 12. Reliability plot of the 48-hr probabilistic forecast of maximum water surface elevation against the  
 463 best-track hindcast for the three hurricane events (Irma, Florence, Laura). The surrogate model results are  
 464 compared to direct empirical evaluation of the Korobov sequence training set used to generate the surrogate  
 465 model, as well as to direct empirical evaluation of the randomly generated 128-member validation set.

489 was found to provide generally more reliable probabilistic predictions than the direct empirical  
 490 evaluation of the training set – the added value of the surrogate model – except for Laura at moderate-  
 491 high probabilities of exceedance. Moreover, the surrogate model offers additional benefits: (1)  
 492 Can be used to rapidly predict the water surface elevations and inundation for any new storm  
 493 perturbation; (2) Provides robust global sensitivity information, and; (3) The water surface elevation  
 494 extrapolation step can be used to purposely bias surrogate inundation prediction low or high, as  
 495 desired.

496 For determining the adequate number of training samples from the Korobov sequence, our  
 497 results show that the surrogate trained on a smaller number of samples (19 here) can provide  
 498 similarly reliable probabilistic predictions as the surrogate trained on more samples (39 or 59  
 499 here). Nevertheless, other results show that the surrogate from 59 samples is more accurate  
 500 ( $RMSE_{90} \approx 1$  ft) than the surrogate from 19 samples ( $RMSE_{90} \approx 2$  ft) when compared to the model





473 FIG. 13. Predicted inundation area of the 48-hr probabilistic forecast as a fraction of the inundation area of  
 474 the best-track hindcast for the three hurricane events (Irma, Florence, Laura). The surrogate model results are  
 475 compared to direct empirical evaluation of the Korobov sequence training set used to generate the surrogate  
 476 model, as well as to direct empirical evaluation of the randomly generated 128-member validation set.

501 simulation validation set. We therefore suggest that if only the probabilistic prediction is required,  
 502 and in a short time frame, that the user could select to train the surrogate using a smaller sample size  
 503 (e.g., 19 which covers 90% of the distribution of  $\lambda$ ). This is in line with previous related research  
 504 that suggested 27 samples was sufficient for this purpose (Davis et al. 2010), and with Kyprioti  
 505 et al. (2021a) who also demonstrates the efficiency of low-discrepancy sequences. Whereas, if the  
 506 user is interested in predicting the water surface elevations for a given perturbation of the storm  
 507 we recommend a higher number of training samples (e.g., 59) be used to produce a more accurate  
 508 surrogate model. It is also possible that for forecasts closer to landfall, fewer training samples will  
 509 be required. Note that the user is free to choose any number of samples, and is not restricted to 19,  
 510 39, and 59 used here.

511 In this study we used PCs due to their simplicity and usefulness in treating uncertainties. The  
 512 use of penalized regression with cross-validation (we recommend Elastic Net with LOO cross-

513 validation) was able to robustly fit the PCs across both the training and validation set, whereas  
514 OLS overfits the training set. Furthermore, it was found that the head loss water extrapolation  
515 technique for filling in dry mesh points was critical to fitting PCs accurately to overland areas.  
516 We also tried extrapolation without head loss but this resulted in poor estimation of inundation  
517 onset (generally has too many false wet predictions). Other studies without head loss extrapolation  
518 (Lee et al. 2021; Plumlee et al. 2021) used CNN and GP machine learning methods that have  
519 more degrees of freedom than PCs, which may help to hide this deficiency. Whether or not this  
520 is viewed positively or negatively, the use of PCs does lead to strong knowledge of the effect of  
521 inputs on surrogate performance. This and the fact that PCs allow for exact extraction of variance-  
522 based sensitivity indices without additional sampling highlights their usefulness for understanding  
523 uncertainty. Nevertheless, future work could explore whether ANN/CNN or GPs can improve  
524 surrogate model accuracy and reliability in our framework.

525 In addition to the accuracy of the surrogate model, reliability is also dependent on the hurricane  
526 perturbation methodology. In this study we followed NHC P-Surge methodology that utilizes  
527 historical statistics of forecast errors. Future work may consider how to perturb hurricane tracks  
528 in a less self-similar fashion (see Fig. 2) and with consideration of the current storm dynamics. In  
529 addition, we validated the reliability against the model simulation of the best-track hindcast, but  
530 real-world observation validation should be assessed in future. This may require the use of more  
531 sophisticated hurricane vortex models to better capture the (potentially asymmetric) wind structure  
532 and storm size. As noted, Laura was the most sensitive to  $R_{max}$  and had the lowest reliability.

533 Though this study focused only on the spatially varying maximum water surface elevation,  
534 the KL-PC methodology can be generally applied to a spatio-temporal surrogate construction, to  
535 account for the temporal evolution of the water surface elevations and hence predict the timing  
536 of the peak flood. Use of log-space surrogate construction to preserve surrogate model positivity  
537 when considering water level time series might be more useful than found in this study. Here,  
538 when pseudo-negative water depths were introduced into the training set from the maximum water  
539 surface elevation extrapolation, the log-space surrogate construction was found to be deleterious  
540 instead of beneficial. This is in contrast to Plumlee et al. (2021) who found the log-transform  
541 necessary for use with GPs.

## 542 **5. Conclusions**

543 A methodology for efficient ensemble perturbation of hurricane wind forcing forecasts and un-  
544 certainty quantification of the resultant simulated coastal flooding has been presented. Probabilistic  
545 prediction results based on the 48-hr forecast prior to landfall for three historical hurricanes are  
546 promising as compared to model simulations of the best-track hindcast. The methodology has  
547 been implemented into a general python framework that can be extended to develop new hurri-  
548 cane perturbation methodologies, use more sophisticated hurricane vortex models, and facilitate  
549 perturbations to parameters in the hydrodynamic model such as bottom roughness.

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566 *Data availability statement.* Data and scripts related to this manuscript are available from Pringle  
567 et al. (2022). EnsemblePerturbation is hosted on the NOAA Office of Coast Survey - Modeling  
568 Github website: <https://github.com/noaa-ocs-modeling/EnsemblePerturbation>, with  
569 “readthedocs” documentation linked therein.

**NHC Historical Forecast Error Tables**

TABLE A1. Mean absolute forecast error: cross-track [nm]

Initial $V_{max}$ (VT=0)			
VT [hr]	< 50 kt	50-95 kt	> 95 kt
0	4.98	2.89	1.85
12	16.16	11.58	7.79
24	23.10	16.83	12.68
36	28.95	21.10	17.92
48	38.03	27.76	25.01
72	56.88	47.51	40.48
96	92.95	68.61	60.69
120	119.67	103.45	79.98

VT: forecast validation time, nm: nautical mile, kt: knot

TABLE A2. Mean absolute forecast error: along-track [nm]

Initial $V_{max}$ (VT=0)			
VT [hr]	< 50 kt	50-95 kt	> 95 kt
0	6.33	3.68	2.35
12	17.77	12.74	8.57
24	26.66	19.43	14.64
36	37.75	27.51	23.36
48	51.07	37.28	33.59
72	69.22	57.82	49.26
96	108.59	80.15	70.90
120	125.01	108.07	83.55

VT: forecast validation time, nm: nautical mile, kt: knot

TABLE A3. Mean absolute forecast error:  $V_{max}$  [kt]

VT [hr]	Initial $V_{max}$ (VT=0) [kt]		
	< 50	50-95	> 95
0	1.45	2.26	2.80
12	4.01	5.75	7.94
24	6.17	8.54	11.53
36	8.42	9.97	13.27
48	10.46	11.28	12.66
72	14.28	13.11	13.41
96	18.26	13.46	13.46
120	19.91	12.62	13.55

VT: forecast validation time, kt: knot.

TABLE A4. Upper and lower bound forecast errors:  $R_{max}$  [sm]

VT [hr]	Initial $R_{max}$ (VT=0) [sm]				
	< 15	15-25	25-35	35-45	> 45
0	[0.00,0.00]	[0.00,0.00]	[0.00,0.00]	[0.00,0.00]	[0.00,0.00]
12	[-17.15,2.47]	[-13.29,5.74]	[-11.26,10.56]	[-14.82,18.24]	[-22.40,25.43]
24	[-23.55,2.31]	[-18.16,9.45]	[-17.93,13.31]	[-12.13,21.01]	[-18.04,34.39]
36	[-24.90,4.20]	[-25.18,9.24]	[-14.88,17.36]	[-11.19,24.89]	[-1.08,43.22]
48	[-30.57,3.64]	[-29.75,9.80]	[-13.36,18.98]	[-8.47,31.64]	[8.46,43.78]
60	[-37.83,1.33]	[-27.25,10.07]	[-13.70,19.29]	[-6.35,31.09]	[8.18,43.14]
72	[-45.11,-0.99]	[-24.75,10.35]	[-14.04,19.60]	[-4.24,30.54]	[7.93,42.51]
96	[-55.26,-3.72]	[-29.71,13.94]	[-11.43,19.67]	[0.37,30.46]	[2.49,38.55]
120	[-61.26,-9.56]	[-35.46,11.77]	[-11.71,19.62]	[-0.84,32.59]	[3.19,40.56]

VT: forecast validation time, sm: US statute mile.

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