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5 **Visualizing Best and Worst Case Scenarios in**
6 **Joint, Constrained, and Time-Dependent Inversions I:**
7 **Null-Space Transfer and Image-Space Contradictions**

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10 **SUMMARY**

11
12 Because geophysical inversion is used in many vital societal applications, it is unfortunate that
13 some aspects of inverse methods are so abstract. The difficulty of identifying fundamental be-
14 haviors is exacerbated when investigating large non-linear problems which combine multiple
15 datasets into a single model, or which produce multiple models with constraints between them.
16 In this first of multiple papers, we investigate and visualize fundamental behaviors of these ab-
17 stract methods beyond what has been described previously by using simple problems. Instead
18 of using the common resolution description, we use the concepts of the Null Space and Im-
19 age Space. After providing readers with an intuitive sense of the behaviors of simpler inverse
20 methods, we investigate cases of Joint, Constrained, and Time-Dependent inversion without
21 errors, before moving on to the influence of errors. We then extract the fundamental behaviors
22 of these complex methods from the presented best and worst cases. These new insights allow
23 us to propose four avenues to improve inversion results (including two novel methods), which
24 we present with similar simple problems. Overall, we show the benefits of producing multiple
25 estimated models using constraints over combining the inverse problems into a single model,
26 and, the benefit of visualizing simple problems to uncover deep insights into the fundamentals
27 of our everyday methods.

28 **Key words:** Joint inversion – Constrained inversion – Null Space – Time-Dependent Tomog-
29 raphy – asynchronous data – Optimized Experimental Design

30 1 INTRODUCTION

31 Geophysical inversion is used in many societal applications. These include geothermal energy
32 (e.g., Jousset et al. 2011; Rawlinson et al. 2012; Soyer et al. 2018), groundwater remediation (e.g.,
33 Bloem et al. 2020), tunnelling and road building (e.g., Hellman et al. 2017), land-slide risk as-
34 sessment and mitigation (e.g., Malehmir et al. 2016), permafrost investigations (e.g., Wagner et al.
35 2019), aquifer characterization (e.g., Doetsch et al. 2010), subduction zone characterization (e.g.,
36 Wagner et al. 2007), nuclear site characterization (e.g., Tso 2019), mining exploration (e.g., Astic
37 et al. 2021; Horo et al. 2021), and volcanic (e.g., Paulatto et al. 2019) and tectonic hazard assess-
38 ment and mitigation (e.g., Hardt & Scherbaum 1994; Kraft et al. 2013; Rawlinson et al. 2012). In
39 these applications, geophysical methods are often used by themselves for individual campaigns, or
40 compared with results from other methods (known as method integration, e.g. Jousset et al. 2011;
41 Malehmir et al. 2016). Improved results can further be obtained using synergies between meth-
42 ods through Joint or Constrained inversion. This can be done using multiple surveys of the same
43 geophysical method (e.g. Julian & Foulger 2010; Horo et al. 2021), and by combining different
44 geophysical methods (Vozoff & Jupp 1975). The latter greatly complicates the interpretation, as
45 the different methods often have different resolutions in time and space, as well as sensitivities
46 to different subsurface properties. Nevertheless, Joint and Constrained inversion have been used
47 to great effect while making the following survey combinations: Seismic Refraction and Electric-
48 al Resistance Tomography (ERT; e.g. Doetsch et al. 2010; Hellman et al. 2017; Wagner et al.
49 2019), ERT and Ground-Penetrating Radar (e.g. Linde et al. 2006), Passive Seismics and Active
50 Seismics (e.g. Wagner et al. 2007), Seismics and Gravimetry (e.g. Paulatto et al. 2019), Receiver
51 Functions and Surface-Wave Dispersion (e.g. Julia et al. 2000), Gravity and Magnetics (e.g. Zhou
52 et al. 2015), Magnetotelluric (MT) and Radio Magnetotelluric (e.g. Commer & Newman 2009),

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53 seismic refraction and MT (e.g. Gallardo & Meju 2007), MT and Local Earthquake Tomography
54 (e.g. Demirci et al. 2018). Additionally, frameworks have been and are being developed to com-
55 bine a larger set of geophysical methods into Joint inversions (e.g. Moorkamp et al. 2011; Rücker
56 et al. 2017), or geologically consistent inversions (e.g. Soyer et al. 2018; de la Varga et al. 2019;
57 Astic et al. 2021).

58 We highlight Time-Dependent inversion in the title, as this method is applied to the same vari-
59 able(s) using the same solver, and needs to identify whether differences in the estimated models
60 are significant beyond the known artificial sources of such differences (Hobé et al. 2021). In com-
61 parison, the other methods tend to assume an unchanging subsurface and are applied between
62 variables, and/or between methods.

63 The complex nature of the individual methods and the in-depth knowledge required for using
64 them individually often makes it hard to get a grip on the fundamental processes and influences
65 in Joint and Constrained inversion. Many synthetic investigations have been developed and are
66 used regularly to gain understanding of the accuracy and resolution of these combined methods.
67 These include checkerboard tests, inverting data subsets, and hypothesis tests (inverting a specific
68 synthetic model, see e.g. Koulakov et al. 2013). For time-dependent tomography, feature robust-
69 ness and the level of artificial differences can be identified using a baseline reconstruction (Hobé
70 et al. 2021), or by using a method for "ground truthing" (Bloem et al. 2020). Another method for
71 Time-Dependent Tomography we will investigate in detail is inter-model minimization (Julian &
72 Foulger 2010), which we refer to as "epoch-damping".

73 In this paper, we present simple inversion problems to visualize fundamental behaviors in geo-
74 physical inversion to ultimately visualize the best and worst case scenarios in Joint, Constrained,
75 and Time-Dependent inversion. The results in this paper are divided into two parts. In Part I, we
76 investigate cases without errors. The Joint, Constrained, and Time-Dependent inversion results are
77 quite complex. Therefore, this part starts with an incremental set of simpler investigations. These
78 investigations both build up towards the complex examples, and help develop the underlying be-
79 haviors, which help explain their best and worst cases. In Part II, we investigate the influence of
80 errors. This part relies heavily on the specifics introduced in Part I.

81 In the discussion, we provide four avenues for improving inversion results, using the funda-
82 mental behaviors identified in the results. A non-linear case study investigating these phenomena
83 will be presented in a following paper.

84 **2 METHODS**

85 All the following examples have been produced in Matlab (MATLAB 2021) using Singular Value
86 Decomposition (SVD) and tested using Conjugate Gradients (e.g., Menke 2018). Because of the
87 simplicity of the examples, the results are equivalent within numerical precision.

88 **3 PART I: CASES WITHOUT ERRORS**

89 Before complicating the picture with the inclusion of errors, this Section looks at inversion cases
90 without errors. This Section is divided into two parts. In the first part we establish some vital
91 concepts and vocabulary with a very simple example. We then dive into the best and worst cases
92 for Joint, Constrained, and Time-Dependent inversion for cases without errors in the second part.
93 Experienced readers may skip most of this first part. For a clear understanding of the later Sec-
94 tions, however, we do recommend having a look at the Geometries Section (Section 3.1.2) and the
95 concepts of Null-Space transfer and Image-Space contradictions (Section 3.1.6).

96 **3.1 Build Up Using Simpler Investigations**

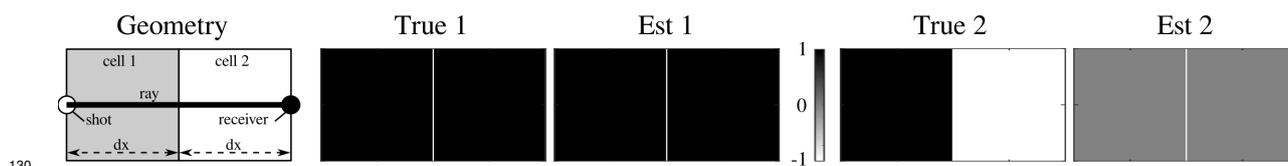
97 In this Section, we present simpler investigations, using vital visuals and vocabulary, to prepare
98 for the complex interactions seen in Joint, Constrained, and Time-Dependent inversions. This Sec-
99 tion is meant as a visual tutorial to help the reader (especially those unfamiliar with some aspects
100 of geophysical inversion) get an intuitive sense of inversion behavior, which we will rely upon
101 when explaining the more complex cases. We expect experienced readers to also benefit from
102 this visual representation, as we learned a lot ourselves from investigating these methods in this
103 way. First, we introduce the Image Space and Null Space using the simplest possible example in
104 Section 3.1.1. Next, we show the simple geometries and corresponding model-space vectors used
105 throughout most of this paper (Section 3.1.2). We also show how changes in the geometry lead to a

106 reorientation of the model-space vectors. After adding the data-space vectors of one of the geome-
 107 tries, we have all the components required to give an intuitive explanation of how SVD behaves
 108 (Section 3.1.3). We then use smoothing to show how constraints change the data-space vectors
 109 and how this changes which model-space vectors end up in the result (Section 3.1.4). Here, we
 110 describe the fundamental workings of the trade-off curve between roughness and data misfit using
 111 the described behaviors. In the last part of this build up (Section 3.1.6), we introduce the concepts
 112 of "Null-Space Transfer" and "Image-Space Contradictions" using an oversimplified non-linear
 113 investigation. Both these concepts help us explain which components of the initial model remain
 114 in the final results, as well as describing a downside of using incorrect geometries in non-linear
 115 inversion when iterating towards a solution.

116 3.1.1 The Simplest Example

117 Fig. 1 shows the simplest possible geometry for a geophysical inversion involving traveltimes. This
 118 is an under-determined problem, as there are two unknowns (2 cells) and one data point (1 ray = 1
 119 traveltimes). True 1 and True 2 are the two "model-space vectors" (MSVs) and could represent one
 120 possible true model each (except that negative slowness values are nonphysical). Every possible
 121 true model for this two-celled example can be reproduced through a linear combination of these
 122 orthogonal MSVs. However, only multiples of True 1 can be reproduced using inversion with the
 123 geometry at hand. For True 2, the traveltimes of the raypath parts cancel each other out. (Given
 124 $v_{21} = 1$ km/s, and $v_{22} = -1$ km/s, $tt_2 = dx/v_{21} + dx/v_{22} = 0$ s.) Therefore, this model component
 125 does not affect the data. The set of MSVs that do not affect the data are said to be in the "Null
 126 Space" (NS). In mathematics, the complement of the Null Space is the "image" (e.g. p. 6 Sharipov
 127 2004), which we will call the Image Space (IS) here ($A = IS(A) + NS(A)$). Thus, the Image Space
 128 (not a commonly used term in geophysics but vital here) is the set of all the MSVs that affect the
 129 data.

135 Consequently, only linear combinations of MSVs in the IS can be reproduced. Therefore, this
 136 geometry can only discern the average of the two cells, and never their differences, as there is
 137 insufficient data to argue for further model complications (c.f. Occam's razor). Visually, MSVs in



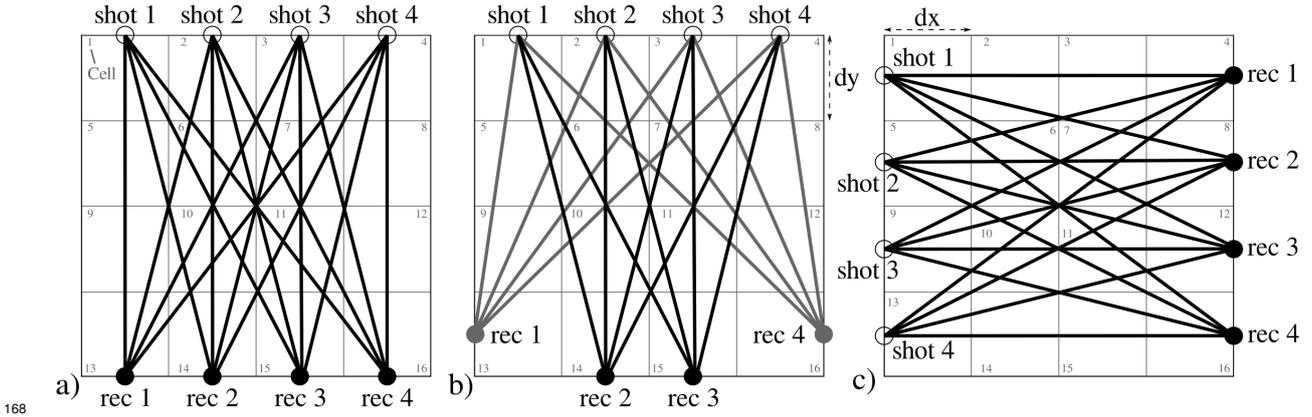
130
131 **Figure 1.** Simplest possible case of an under-determined problem using traveltimes. The two model vec-
132 tors (True), which combined can produce any possible model using this geometry, and their least-squares
133 inversion using SVD (Est). True 1 lies in the Image Space, and can therefore be reproduced up to numerical
134 precision, whereas True 2 lies in the Null Space. Its estimated model is zero up to numerical precision.

138 the NS produce model components that are zero up to numerical precision (their corresponding
139 singular values are zero, or near zero). Adding Est 1 and Est 2 is thus basically equivalent to Est
140 1. Practically, MSVs in the NS are excluded from the inversions using e.g. truncated-SVD, which
141 leads to the same result.

142 Although our examples in this paper use traveltimes, the same reasoning and insights apply
143 either directly or indirectly to other geophysical methods. To demonstrate this on the simplest
144 level, we may parameterize a gravity experiment using a single measurement and two cells: a top
145 layer above a halfspace. Because there is only one measurement, we would again only be able to
146 produce one density for the entire model. For ERT, we could get a single measurement using e.g. a
147 Schlumberger array. A 2 cell parameterization would again produce the same two MSVs, and we
148 would again only be able to estimate the one uniform resistivity throughout the model.

149 3.1.2 Geometries

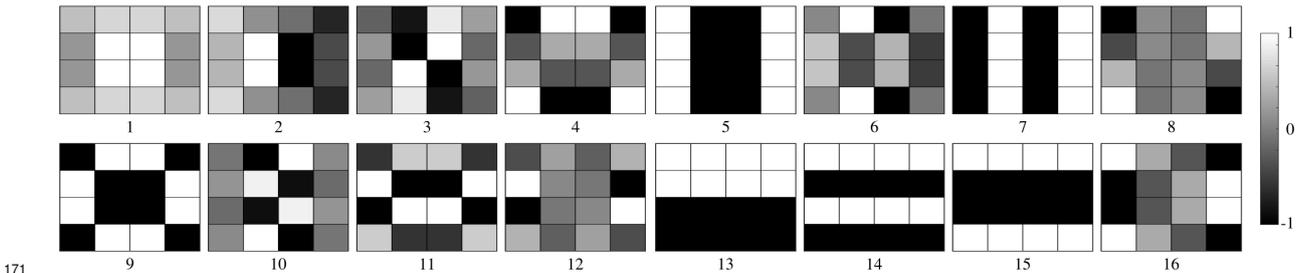
150 In this paper, we use the geometry in Fig. 2a to show the fundamental behaviors of a range of
151 inversion methods. This geometry (modified from Lévêque et al. 1993) was chosen for it's visual
152 clarity and again presents a seismic travelttime problem. Fig. 2b shows a slight adjustment to the
153 Leveque geometry (compare rec 1 and rec 4). This adjusted geometry is used in the non-linear
154 investigation together with the Leveque geometry to show Null-Space effects arising in non-linear
155 inversion (Section 3.1.6). We chose the horizontal geometry (Fig. 2c) to provide a maximum dif-
156 ference to the vertical geometry when using Joint, Constrained, and Time-Dependent inversion.
157 All cases assume straight rays between the shots and receivers ($dx = dy = 1$ km). The orthog-
158 onal MSVs (Figure 3) are then produced using SVD of the geometry's G-kernel. Similar to the
159 2-cell case (Fig. 1) these MSVs can be combined to produce every possible model for this 4x4



168 **Figure 2.** Model geometries. 16 cells with **a)** vertical setup (modified after Lévêque et al. 1993), **b)** adjusted
 169 vertical setup, and **c)** horizontal setup.
 170

160 discretization. These MSVs will be used as our "true" models in this paper. Note that any 4x4 cell
 161 parameterization will produce the same set of MSVs reoriented to accommodate the geophysical
 162 problem at hand. In other words, the MSVs produced by applying SVD on a 4x4 cell parame-
 163 terization for any geophysical method, will each be linear combinations of the MSVs in Fig. 3.
 164 This is also true of 4x4 nodal parameterizations, though the resulting models would look different
 165 due to the interpolation between nodes, the values on the nodes will be linear combinations of
 166 the MSVs shown here. **Thus, the main insights in this paper should be reproducible with any**
 167 **other geophysical method (or combination of methods) that employs inversion.**

172 Table 1 compares the reproducibility of these MSVs between the three geometries. In essence,
 173 this table describes what happens if we use a given MSV as the true model to produce synthetic
 174 data. For the vertical geometry (Fig. 2a), the first 12 vectors lie in the Image Space (IS), whereas
 175 the last four lie in the Null Space (NS). These latter four vectors lie completely in the IS for the
 176 horizontal geometry (Fig. 2c). Instead, vectors 4, 5, and 7 lie in the NS of the horizontal geometry.
 177 This makes sense, as they are rotated versions of vectors in the vertical geometry's NS (16, 15, and



177 **Figure 3.** Stretched Model-Space vectors of the vertical geometry (modified from Lévêque et al., 1993).

Table 1. Model-space vector (**MSV**) comparison for the three geometries (**Geom.**) in Fig. 2, using the MSVs of the vertical geometry. MSVs lie in the Image Space (**IS**), Null Space (**NS**), or both (**B**). *98 % NS

Geom. \ MSVs	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16
Vertical	IS	NS	NS	NS	NS											
Horizontal	B	B	IS	NS	NS	B	NS	IS	IS	IS	B	IS	IS	IS	IS	IS
Vertical adjusted	B	IS	IS	B	B	IS	IS	IS	B	IS	B	IS	B*	B*	B*	IS

178 14, respectively). In comparison, models 1, 2, 6, and 11 have both IS and NS components using
 179 the horizontal geometry. For the adjusted vertical geometry, all vectors lie either in the IS or both
 180 in the IS and NS. However, models 13, 14, and 15 lie almost completely in the NS.

181 The comparisons in Section 3.2 require the orthogonal behavior seen in Table 1 for vectors
 182 4, 5, 7, and 13-16. The SVD algorithm does not produce this "pure" orthogonality in the MSV.
 183 Therefore, slight adjustments were made to these components to remove unwanted contamination
 184 and to properly represent vectors 13-16. These vectors were additionally stretched for plotting
 185 purposes.

186 One important take away from Table 1 is the following: A change in the geometry will change
 187 which MSVs lie within either the IS or the NS. This is important for Time-dependent inversion, as
 188 changes in geometry can thus produce large differences for the same true subsurface. The actual
 189 representation of these MSVs after applying SVD on a new geometry can also be linear combi-
 190 nations of the MSVs in the original geometry. Changing between MSV's for different geometries
 191 thus merely represent a re-orientation of the original MSV's.

192 Fig. 4 shows the data-space vectors (DSVs) corresponding to the MSVs in Fig. 3 along with
 193 the singular values. These orthogonal vectors can be seen as data components of the vertical prob-
 194 lem. Using linear combination, these 16 DSVs can produce every possible data vector for the 16
 195 rays. These DSV also each correspond to the data which each MSV would produce. As the DSV
 196 themselves are normalized, the singular value is used to scale the DSVs to obtain the synthetic data
 197 of each MSV. Vectors 13-16 are clearly in the NS as their singular values are close to numerical
 198 precision.

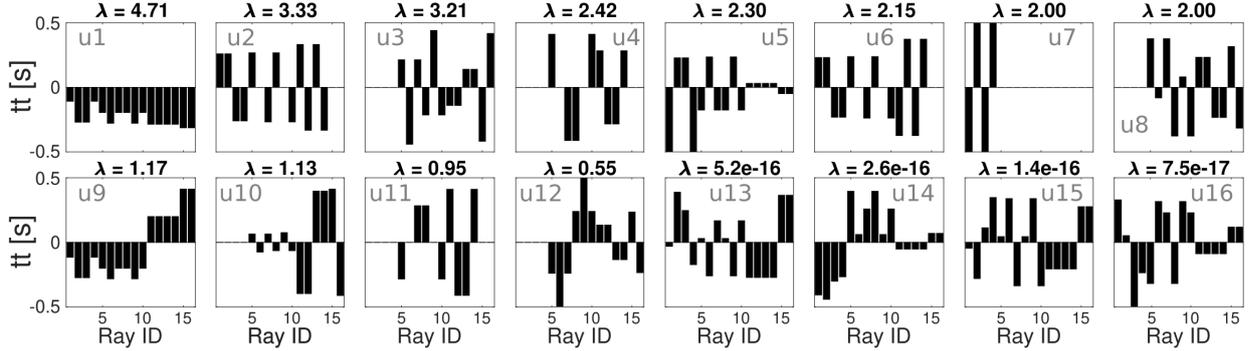


Figure 4. Data-Space vectors corresponding to the Model-Space vectors in Fig. 3. Values at the top of each subplot correspond to that vector's Singular Value.

199 3.1.3 SVD Behavior

200 Now that we have all of the required aspects involved in SVD inversion, we will give an intuitive
 201 explanation of how SVD translates data into the least-squares estimate.

202 Eq. 1 shows how the forward kernel, G , (which holds the ray lengths for each cell in our
 203 problems) can be decomposed into three matrices, and Eq. 2 shows how to arrange these matrices
 204 into the least-squares solution (e.g., Menke 2018):

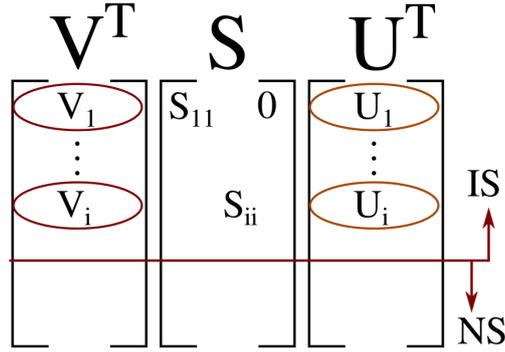
$$G = USV^T \quad (1)$$

$$m^{Est} = VS^{-1}U^T d^{obs} \quad (2)$$

205 Here, m^{Est} , is the estimated model, and d^{obs} , the observed data. U , S , and V are explained
 206 in Fig. 5. We refer the reader to e.g. Menke (2018) for a mathematical derivation of how SVD
 207 decomposes the problem into the individual matrices.

215 Let us first look at $U^T d^{obs}$ in Eq. 2. Here, a dot product is formed between each DSV and the
 216 data. Thus, we check the length of the data in the direction of each DSV. If this length is non-
 217 zero, the corresponding MSV will be activated. How much of each MSV is used in the final result
 218 is determined by using a weight. This weight is obtained using the singular values as follows:
 219 $S^{-1}U^T d^{obs}$. The final estimated model is then obtained by multiplying each MSV with its weight
 220 and adding all the resulting model components.

221 Because a singular value of $1e-16$ would produce unrealistically large models, these vectors



208

209 **Figure 5.** Singular Value Decomposition decomposes the G-kernel into three matrices, V, S, and U. The
 210 columns of V hold the model-space vectors, and the columns of U hold the data-space vectors. S is a
 211 diagonal matrix which holds the singular values. In this image, the three matrices are oriented to align the
 212 Null Space and what we call the Image Space. Note that this orientation does not correspond to either the
 213 decomposition of G (Eq. 1), nor the least-squares solution (Eq. 2). The Null Space part of the model space
 214 can also be larger or smaller than the Null Space part of the data space.

222 are usually removed from the equation using truncated-SVD (e.g, Menke 2018). The behavior of
 223 SVD can be altered, however, to make use of the Null-Space entries. One way of doing so is to
 224 provide constraints, which we will look at next.

225 3.1.4 Smoothing constraints

Smoothing constraints diminish differences between adjacent cells. These constraints are often used in under-determined (or mixed determined) problems to overcome the ill-posedness of the problem. The problem is thus said to be regularized, which makes the problem behave better during inversion. This also means that we change the question asked of the data. In this Section, we present how such a change in the question fundamentally changes the entries in the U, S, and V matrices. Insights into this allows us to explain how an increase in the smoothing constraints leads to an increase in the data misfit. First we will describe the extension of the kernel and data vector required for smoothing. Implemented as a soft constraint, the smoothing operator multiplied with the model vector equals zero:

$$\begin{bmatrix} G \\ \alpha D \end{bmatrix} m = \begin{pmatrix} d^{obs} \\ \vec{0} \end{pmatrix} \quad (3)$$

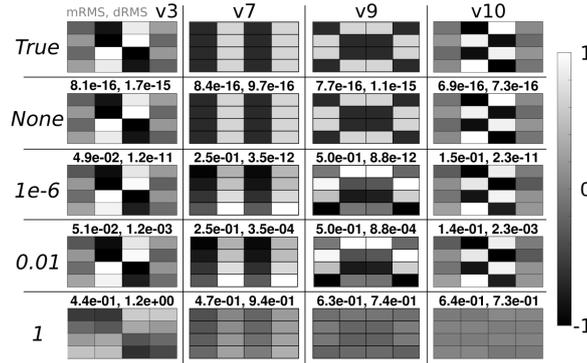


Figure 6. Examples of inversion using the vertical geometry with different levels of smoothing (without, 1e-6, 0.01, and 1) and model-space vectors v3, v7, v9 and v10 as the true models. Values correspond to model misfit RMS (mRMS) and data misfit RMS (dRMS). We note that the model with the highest singular value (v3) is the least affected by the smoothing.

where, D , is the Laplacian operator

$$D_{i,j} = 4m_{i,j} - m_{i-1,j-1} - m_{i-1,j+1} - m_{i+1,j-1} - m_{i+1,j+1} \quad (4)$$

and, α , is the smoothing factor. The Laplacian thus adds one row for every model cell, where index i and j in Eq. 4 are the model cell indices in the x - and y -directions, respectively. Thus, α controls how strongly this regularization is emphasised relative to the data equations. It is important to note here, using the standpoint of image processing, that providing the Laplacian values of the true MSVs (the right side of Eq. 3) would allow us to perfectly reconstruct any model for this geometry. Because these values are not known, we force the solution to be smooth (i.e. the zero entries instead). We visualize the behavior of smoothing regularization with the following synthetic reconstruction tests.

We use the MSVs of the unregularized vertical problem (Fig. 3) as true models, with which we produce synthetic traveltimes. These traveltimes are then inverted using different levels of smoothing (none, 1e-6, 0.01, and 1). Four of the resulting estimated models are shown together with the true models in Fig. 6.

These results show that smoothing constraints come at a cost of both the model misfit (mRMS) and the data misfit (dRMS). The model recovery is poorest for the strongest regularization ($\alpha = 1$). The recovery at high α values diminishes with decreasing singular value (see Fig. 4 for comparison). Fig. 6 also shows that some cell values may increase or decrease, to minimize the

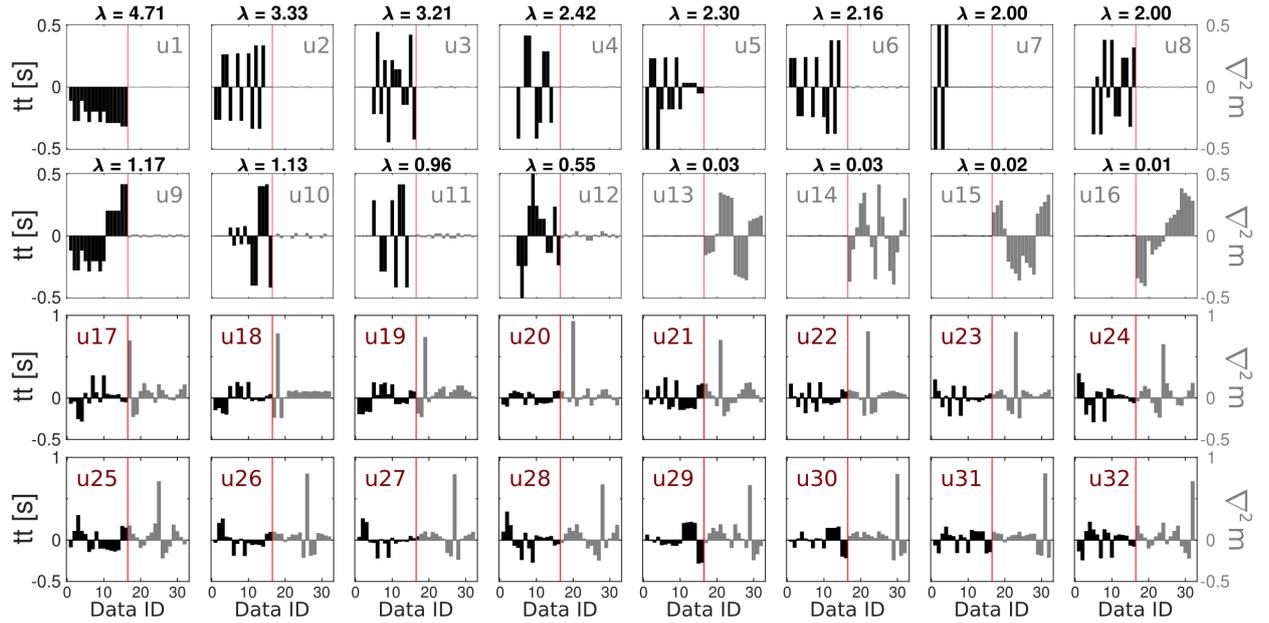


Figure 7. Data-Space vectors (e.g. u1) of the smoothed vertical problem ($\alpha = 0.01$). Values to the left of the red line are normalized traveltimes. Values to the right of the red line are normalized Laplacian values of each cell (i.e. larger values mean larger differences between adjacent cells). These latter values scale with the smoothing factor α . λ values at the top are each vector’s singular value. u17-u32 do not have corresponding singular values and thus lie in the Null Space.

242 impact on dRMS as the estimated models are smoothed out (e.g. $\alpha = 0.01$ for v7 and v9). Next,
 243 we will present both the data-space vectors (Fig. 7) and the model weights associated with these
 244 synthetic reconstruction tests (Fig. 8), to provide visual clarity of how regularization allows the
 245 inversion to contradict the traveltimes information.

246 The added rows in Eq. 3 show up in the data space as additional data-space vectors (Fig. 7):
 247 one for each constraint equation in D . Each DSV is also extended beyond the traveltimes (see Fig.
 248 4 for comparison). When looking at the new DSVs (bottom two rows in Fig. 7) there are a few
 249 things that stand out. Firstly, the Laplacian entries on the right side of the red line (representing
 250 the border between the traveltimes data and the Laplacian values) each show a single large spike
 251 corresponding to the central cell (see Eq. 4). Secondly, these new DSVs have non-zero values for
 252 the traveltimes entries. Similarly, non-zero Laplacian values have been added to u1-u16. Their size,
 253 relative to the traveltimes, scales with the singular values. Here, DSVs u2, u3, u4, u8 and u11 have
 254 ”reversed polarity” relative to the unregularized vertical problem (c.f. Fig. 4). The corresponding
 255 MSVs have changed sign accordingly. We also note that u13-u16 no longer reside in the NS
 256 ($\lambda \geq 0.01$). In this case, with $\alpha = 0.01$, the singular values are still two orders of magnitude

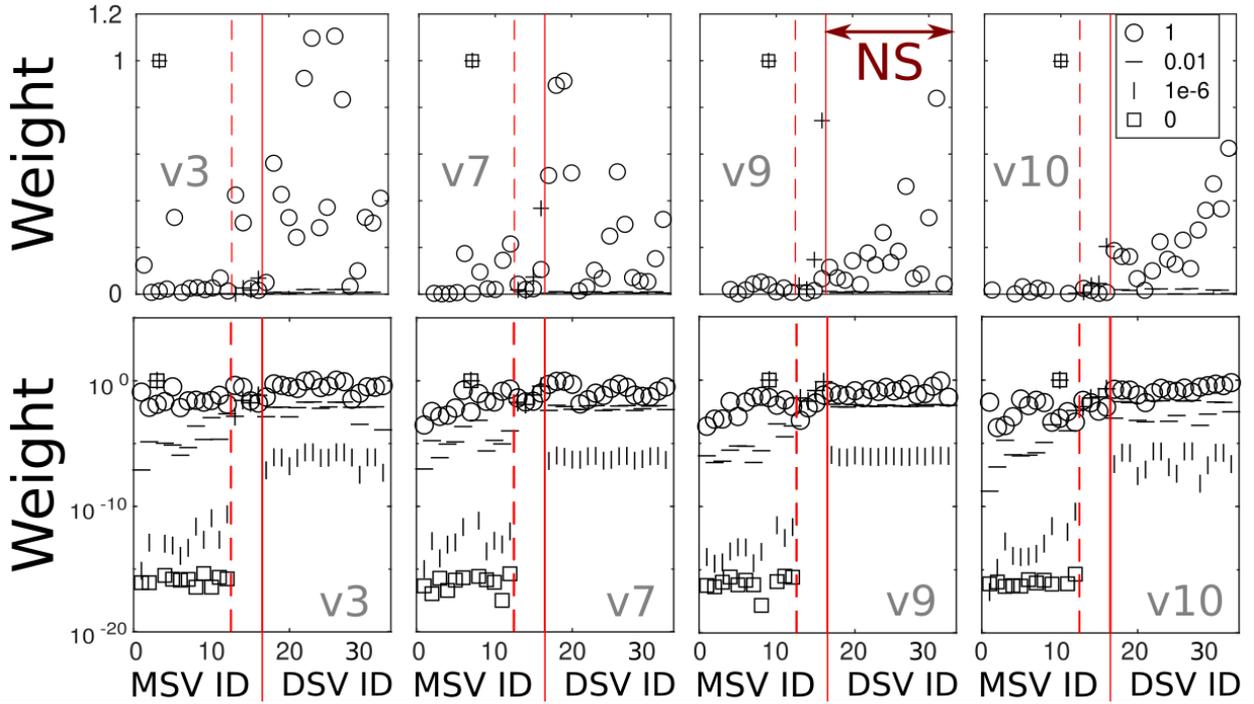


Figure 8. Reconstruction examples at multiple smoothing levels (without, $1e-6$, 0.01 , and 1) with model components of the vertical problem as true models (e.g. $v3$). The top and bottom rows are in linear and log-linear scale for emphasis, with the top row not showing values smaller than $1e-3$ for visual clarity. To the left of the full red line, these plots show the weights for each of the model-space vectors (MSVs) of the smoothed vertical problem, i.e., how much of each MSV is used when inverting for a given true model. Values between the broken and the full red lines correspond to the four MSVs that are in the NS for the unregularized problem. The values to the right of the full red line instead show how much each DSV in the NS is activated.

257 smaller than those for $u1-u12$, but they scale with α . The main contribution of these four DSVs is
 258 fitting the Laplacian and less to the travel time data.

259 The results in Fig. 8 correspond to the four cases in Fig. 6 (see the caption of Fig. 8 for further
 260 details). Because Eq. 3 does not change the parameterization (i.e. the size of m), the smoothed
 261 vertical problem also has 16 MSVs. As there are 32 DSVs, however, there is more opportunity
 262 for the data vector to activate DSVs that do not map onto the estimated model, compared to the
 263 unregularized vertical problem.

264 The results in Figs. 8 show that each reconstruction preferentially activates one MSV when α
 265 does not overpower the traveltime contributions (e.g. $\alpha = 1$). (Results for all cases are presented in
 266 Appendix A.) When looking at the log-linear view, however, we see that all vectors see increased
 267 activation as α increases. Although it is not clear from these images, the weight on the preferen-
 268 tially activated MSV is lowered accordingly (e.g. the weight on MSV 3 for case $v3$ is lowered by
 269 $\sim 1e-5$ at $\alpha = 1e-6$). Especially interesting is the activation of MSVs between the broken

270 red line and the full red line (see e.g. the apparent pluses in case v9). Even small α values thus
 271 activate model components that were in the NS for the unregularized vertical problem.

272 All of the above observations describe consequences of vector re-orientation. Similar to the
 273 vector redistribution when changing the geometry (Table 1), the MSVs of the unregularized verti-
 274 cal problem get redistributed over the new MSVs in the regularized problem. The same occurs for
 275 the DSV, except that the traveltimes components now get redistributed over 32 DSVs instead of 16.
 276 The synthetic data for these reconstructions thus now (partially) activate all DSVs, because each
 277 DSV includes a component aligning with the original DSVs. As α increases, more of each origi-
 278 nal DSV in Fig. 4 is redistributed among all DSVs in the regularized problem (seen as increased
 279 activation of all vectors in Fig. 8). As a consequence, all MSVs of the regularized problem are
 280 included to some degree in our synthetic reconstructions. As these MSVs also include large parts
 281 of the other model components in Fig. 3, their activation incorporates components not included
 282 in the true model. The regularized problem thus allows the inclusion of other model components
 283 and contradiction to the traveltimes by re-orientating the vectors and having data components map
 284 onto the NS of the data.

285 3.1.5 *A closer look at the L-curve*

286 Using these insights and observations we can now explain the behaviors observed in the "L-curve"
 287 commonly used to find a suitable value for α . By plotting dRMS as a function of model roughness
 288 (i.e. inverse of smoothness) this trade-off curve behaves as follows. As α increases, the roughness
 289 decreases rapidly with a small increase in dRMS (the horizontal part of the L). This behavior
 290 switches entirely in the corner of the L, to instead have large increases in dRMS for small further
 291 decreases in roughness (the vertical part of the L). As seen in Fig. 7, the MSVs formally in the NS
 292 (e.g. MSV 16 for v9) get activated strongly at small α values and to a lesser degree when $\alpha = 1$.
 293 When we examine the corresponding vectors u13-u16, we note that the relative amplitude of the
 294 constraint values is much larger than the traveltimes values, while the opposite occurs for u1-u12.
 295 Thus large changes to the estimated models are obtained without incurring penalties in the dRMS
 296 by incorporating those MSVs that only affect the data marginally. As α increases, more of the true

297 model vector (e.g. v_3) is included in all MSVs, leading to more inclusion of model components that
 298 contradict the data. The combined behavior in the two previous sentences explains the horizontal
 299 part of the L-curve.

300 As α increases further, however, all model components get redistributed further, including
 301 those that do not affect the data. (All MSVs in the smoothed problem thus map onto both the
 302 IS and NS of the unregularized problem.) Activating MSVs 13-16 in the smoothed problem thus
 303 incorporates more and more model components that affect the data. The corner of the L signifies
 304 the moment where the desired model vector is diluted so much over the available MSVs that
 305 several DSVs in the IS are activated significantly at the same time by the desired traveltimes, thus
 306 incorporating larger and larger components not included in the true model. This is best seen for v_3
 307 (Fig. 8) for $\alpha = 1$. Again, the misfit itself arises from traveltimes components activating the NS.

308 Although this investigation does not directly address the main goal of identifying the best
 309 and worst case scenarios of Joint, Constrained, and Time-Dependent inversion, the mechanisms at
 310 play are the same. How the IS and NS affect geophysical results is further discussed in the next
 311 Sub-Section and will prove to be vital for our main goal.

312 3.1.6 Non-Linear and Initial Model Investigation

313 To show the influence of changes in the Image-Space and Null-Space in non-linear inversion,
 314 we perform a simplified non-linear investigation. This investigation uses the following steps: 1)
 315 Produce synthetic data for one of the model-space vectors in Fig. 3 using the vertical geometry
 316 (Fig. 2a), which we will call the final geometry here. 2) Choose an initial model. 3) Produce a
 317 model update using the adjusted vertical geometry (Fig. 2b), which we will call the intermediate
 318 geometry here. 4) Produce a final model update using the final geometry. This procedure simulates
 319 the situation in which we do not know the exact ray path locations.

320 Five such simulations are shown in Fig. 9. Sims. 1 to 4 use MSV 1 as the true model, and
 321 simulation 5 uses MSV 2. Both of these lie in the IS of the vertical geometry. Simultaneously,
 322 MSV 2 lies in the IS of the intermediate geometry, whereas MSV 1 lies partly in its NS. The
 323 first simulation use a starting model of zero, which essentially means there is no starting model,

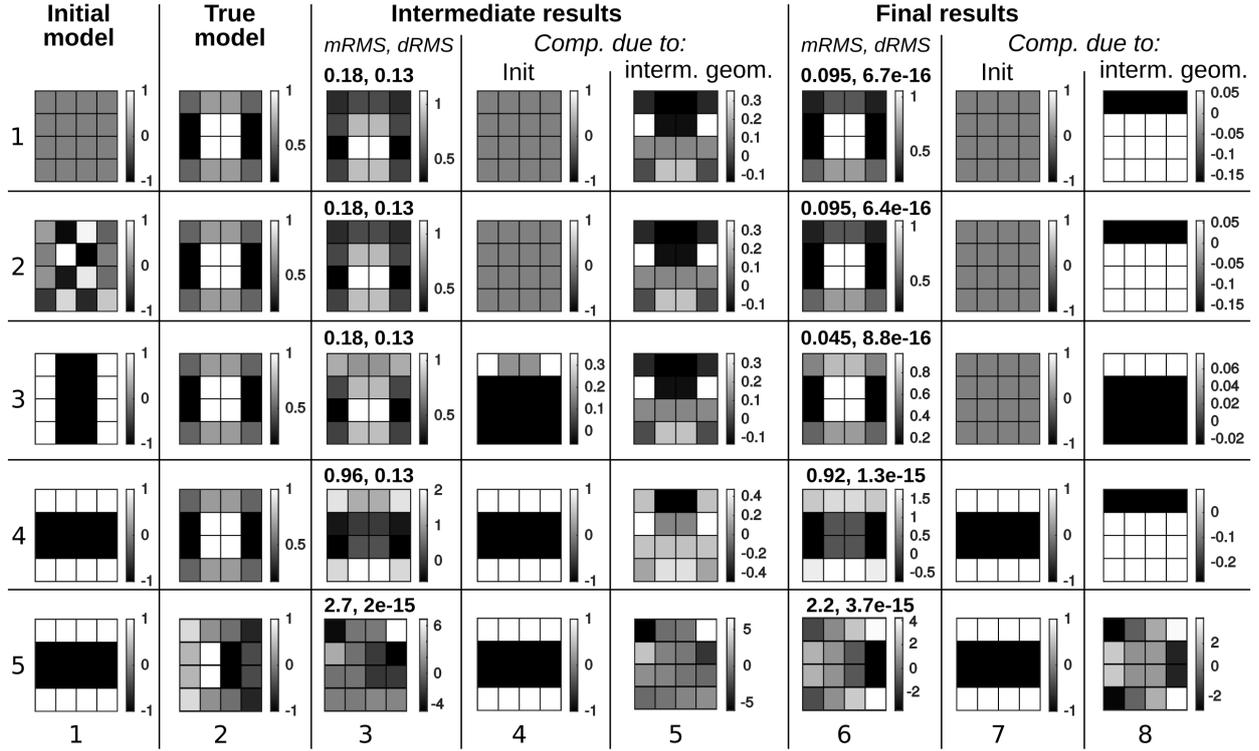
324 and is done to show a result free of its influence. Simulation 2 uses a starting model in the IS of
325 the intermediate geometry, which lies in both the IS and NS for the final geometry. Simulation 3
326 uses a starting model that lies in the IS of the final geometry, and in both the IS and NS of the
327 intermediate geometry. Simulation 4 uses a starting model in the NS of the final geometry, and in
328 both the IS and NS of the intermediate geometry. Simulation 5 uses the same starting model as
329 Simulation 4, and a different true model (see above). Before discussing the results, we explain the
330 contents of each column in Fig. 9.

331 Each row starts with the chosen initial model and the true model (columns 1 and 2). The next
332 three columns (3-5) break down the intermediate result produced using the intermediate geometry.
333 The first column of the intermediate results (column 3) shows the intermediate model. This model
334 can be decomposed into three components: the true model (column 2), the remainder of the initial
335 model (column 4), and a third component produced by using the wrong geometry (column 5).

336 The last three columns (6-8) break down the final results using the final geometry in a similar
337 fashion. The first column of the final results (column 6) shows the final models. The next column
338 (7) shows the model component of the initial model which remains in the final result. The last
339 column (8) shows what remains of the component introduced due to the intermediate geometry in
340 the final result.

353 Simulations 1 and 2 in Fig. 9 have identical intermediate and final models even though their
354 initial models are different. In Simulation 2, the intermediate result is not affected by an initial
355 model, which lies in the IS of that geometry. Because it lies in the IS, the initial model produces
356 synthetic data which contradicts the observations. We call this Image-Space Contradictions. This
357 causes the inversion to resist incorporating this component into the result. From this we conclude
358 that model components of the previous iteration will always be overwritten, if 1) they lie in the
359 current geometry's IS, and 2) the inversion allows it (e.g. smoothing and model-step limitations).

360 The first two simulations also show that the incorrect geometry introduces a model component
361 in the intermediate result, which is mostly removed by the final geometry. The data misfit of the
362 final result (numerical precision) shows that the remaining component lies in the NS of the final
363 geometry. From this we further conclude that any model component of the previous iteration,



341

342 **Figure 9.** Results of the non-linear investigation using four different starting points. From top to bottom,
 343 these are: row 1) zeros everywhere, row 2) a model in the Image Space (IS) of both geometries, row 3) a
 344 model in the IS of the vertical geometry and partially in the Null Space (NS) for the intermediate geometry,
 345 and row 4 & 5) a model in the NS for the vertical geometry and mostly in the NS for the intermediate
 346 geometry. This last case is repeated with a different true model. Shown from left to right: column 1) initial
 347 model, col. 2) true model, col. 3) intermediate result using the intermediate geometry, col. 4) remainder of
 348 the initial model in the intermediate result, col. 5) remaining component of the intermediate result due to
 349 using the intermediate geometry, col. 6) final result using the vertical geometry, col. 7) remainder of the
 350 initial model in the final result, and col 8) the remainder of the component due to the intermediate geometry.
 351 *mRMS* and *dRMS* describe the root-mean-squared misfit of the model and data, respectively. The row
 352 numbers correspond to the individual simulations.

364 that lies in the NS of the current geometry, will always remain. We call this Null-Space Transfer.
 365 Note that this true model can be reproduced up to numerical precision when only using the final
 366 geometry. The component due to the intermediate geometry is therefore an undesired result of
 367 needing to find the correct geometry using non-linear iterations.

368 In simulation 3, the NS component of the starting model remains in the intermediate result. The
 369 final model, however, is unaffected by the initial model. Because the initial model lies completely
 370 in the IS of the final geometry, the remaining component in the intermediate model was removed
 371 from the final model.

372 Simulation 4 shows that its initial model affects both the intermediate and the final results. This
 373 initial model additionally produces a component in the intermediate geometry which lies in the NS

374 of the final geometry. This component is much larger compared to the component in the other rows.
375 Simulation 5 shows even stronger effects, despite the fact that the true model is in the IS of both
376 geometries. The difference here is that the intermediate geometry would produce different data for
377 this IS component, compared to the true geometry. The dRMS of the intermediate model shows,
378 however, that the synthetic data fully maps onto the IS for the intermediate geometry. This causes
379 a larger component to be activated, a large part of which unfortunately lies in the NS of the final
380 geometry.

381 Comparing the different simulations, we see from Simulations 1 to 3 that if the starting model
382 has no component in the NS of the final (and true) geometry, there will be no direct effects of
383 the starting model remaining in the final model. The initial model would thus optimally consist of
384 model components which lie in the IS of the final geometry. Whether this is feasible depends on
385 the geophysical method and the geometry. From Simulations 4 and 5 we see that any components
386 of the starting models that lie in the NS of the final geometry will remain in the final model. An
387 additional complication is how the non-linear inversion develops towards a final model. This devel-
388 opment is affected by the initial model and could introduce intermediate components which remain
389 in the final NS. We thus conclude that any component in the initial or intermediate models of a
390 non-linear inversion, which consistently lies in the NS of the subsequent geometries, will remain
391 in the final model. We call this unwanted form of Null-Space Transfer: Null-Space Contamination.
392 We also note in Simulations 1 to 4 that the data fit is perfect (within numerical precision), despite
393 the models being rather different. Without knowing the answer before hand, it is of course difficult
394 to judge which model is "the best".

The difference between NS Transfer and NS shuttles (e.g., Deal & Nolet 1996; Rowbotham & Pratt 1997) is that NS shuttles are used to get a desired outcome by projecting model features onto the NS and adding them to a given model. This allows practitioners to e.g. add prior knowledge with minimum impact on the data misfit (e.g., Rowbotham & Pratt 1997), quantify uncertainty around the model (e.g., De Wit et al. 2012), and explore the set of acceptable results (e.g., Fichtner & Zunino 2019; Fichtner et al. 2021). NS Transfer, in contrast, is an inherent part of both non-linear inversion, and, as we will see, of Joint, Constrained, and Time-Dependent inversion. The

production of new NS components in the intermediate steps is - to our knowledge - missing from descriptions of non-linear inversion in the literature. The closest description of this phenomenon in literature is the propagation of errors in the following equation (modified from Rawlinson & Spakman 2016):

$$\hat{m} = G^{-g}Gm^{true} - G^{-g}Em^{true} + G^{-g}\epsilon \quad (5)$$

395 With, \hat{m} , and m^{true} , the estimated and true models, G and G^{-g} the (linear) G-kernel and
 396 its generalized inverse, E , a matrix representing the accumulated errors in the G-kernel due to
 397 linearization, and ϵ , an error term combining errors due to linearization, noise, picking, and pa-
 398 rameterization. The three terms in Eq. 5 can be reduced to the following conceptual description.
 399 $G^{-g}Gm^{true}$ corresponds to the mapping of the true model onto the estimated model. $-G^{-g}Em^{true}$
 400 then subtracts the mapping of the linearization errors in the G-kernel compared to the "true" G-
 401 kernel of the true model onto the estimated model. Lastly, $G^{-g}\epsilon$ adds the mapping of all the sources
 402 of errors onto the estimated model.

403 Instead of using the concepts of the IS and NS, Rawlinson & Spakman (2016) use the concept
 404 of resolution to describe this error propagation. Whereas the resolution description tends to work
 405 as a black box which morphs the data and initial model into the estimated model, the concepts of
 406 NS transfer and IS contradictions allow us to clearly describe the fundamental behaviors at each
 407 point of the non-linear process based on the IS and NS components of each geometry.

408 A different issue with non-linear inversion is presented in Simulation 5 in Fig. 9. The interme-
 409 diate geometry reproduces the data up to numerical precision. As this is the wrong geometry, it
 410 produces a very poor intermediate model. Additionally, the NS component remaining in the final
 411 geometry is very large for this case. Because of the low data misfit in the intermediate case, it is
 412 unclear if the final geometry would even be reached, though the model change would produce a
 413 change in geometry and thus in the misfit. Additionally, because this setup is oversimplified, it is
 414 unclear if this issue will arise in true non-linear problems. This issue is commonly dealt with in
 415 software like, e.g. PStomo_eq (e.g., Tryggvason & Linde 2006), by limiting the maximum size of
 416 the model update in a single iteration.

3.2 Joint, Constrained, and Time-Dependent Inversion

Now that we have introduced all the vital concepts, we will move on to visualizing the best and worst case scenarios of Joint, Constrained, and Time-Dependent inversion for cases without errors. First we describe the investigation with which these cases were found. Then we show and discuss cases where Null-Space Transfer dominates. Next we show and discuss cases where IS Contradictions dominate. Lastly, we show and discuss cases where both occur simultaneously. For those readers who skipped ahead, NS Transfer and IS Contradictions are explained in Section 3.1.6.

We will use the following nomenclature to simplify our upcoming discussions. Combined inversion: The group of inversion methods that combine two (or more) datasets, belonging to one (or more) geophysical methods, into one (or more) estimated models. Within combined inversion, there are two categories: those that produce a single model, and those that produce multiple models. We could call these single-model producing methods, and multi-model producing methods. Examples of single-model producing methods include: joint inversion of two ERT arrays (e.g. Horo et al. 2021), and joint inversion of asynchronous data in e.g. local earthquake tomography (e.g. Hobé et al. 2021). Because single-model producing methods is a mouth full, we will use "joint inversion" throughout this text instead (we join two problems to produce a single model). We fully realize that this is only a subset of the methods that are commonly (and inconsistently) described as "joint inversion". Examples of multi-model producing methods include: cross-gradient constraints (e.g. Gallardo & Meju 2004; Gallardo 2007; Tryggvason & Linde 2006; Manukyan et al. 2018), combined inversion using petrophysical relationships (e.g. Haber & Holtzman-Gazit 2013; Wagner et al. 2019), and inter-model minimization (e.g. Julian & Foulger 2010).

3.2.1 *Methods and Setup*

We use the vertical and horizontal geometries (Fig. 2a and c) to investigate fundamental behaviors of combined inversion methods. The true models for this investigation are all permutations of the MSVs in Fig. 3 (one for each geometry). Using these true models, we first produce synthetic data. This data is then inverted for using Singular Value Decomposition and Conjugate Gradients in conjunction with the following three methods at different levels of regularization: 1) Joint

444 inversion, 2) Inter-model minimization (epoch damping), and 3) Equivalent-gradient constraints.
 445 Additionally, the data is inverted for without any regularization, which we call "single inversion".

The following two equations describe the implementation of these inversion schemes:

$$\begin{bmatrix} G_1 \\ \alpha G_2 \end{bmatrix} m = \begin{pmatrix} d_1^{obs} \\ \alpha d_2^{obs} \end{pmatrix} \quad (6)$$

$$\begin{bmatrix} G_1 & \vec{0} \\ \vec{0} & G_2 \\ & & \alpha D \end{bmatrix} m = \begin{pmatrix} d_1^{obs} \\ d_2^{obs} \\ \vec{0} \end{pmatrix} \quad (7)$$

446 Here, G_1 corresponds to the kernel of the vertical geometry, G_2 , is the kernel of the horizontal
 447 geometry, D is the matrix holding the regularization entries described below, and α is the regular-
 448 ization factor (weight). The regularization factors for each of these combined inversion schemes
 449 are: 1, 0.01, 1e-6.

450 The joint-inversion scheme vertically combines the equations of both geometries and adds a
 451 regularization factor to the horizontal geometry equations (G_2 in Eq. 6). Inter-model minimization
 452 can be used in Time-Dependent Tomography (Julian & Foulger 2010) and minimizes model dif-
 453 ferences (i.e. $\alpha * [m1_{ij} - m2_{ij}] = 0$; Eq. 7). Equivalent-gradient constraints (EG) minimizes the
 454 differences in absolute model gradients ($\alpha * [\nabla m1_{ij} - \nabla m2_{ij}] = 0$; Eq. 7).

455 EG is similar to cross-gradient constraints (CG), which minimizes the cross-gradient of the
 456 models (e.g. Gallardo & Meju 2004). We show the results for equivalent gradients here, as this is
 457 a linear operation and CG is fundamentally non-linear. When comparing EG and CG, CG allows
 458 models to have, e.g., negative gradients, or zero gradients, where the other model has a positive
 459 gradient. Thus, the behaviors of CG would go beyond the ones described here for EG, while
 460 simultaneously being complicated by CG's non-linearity.

461 As we have seen in Table 1, MSVs of a given geometry can be completely in the IS, completely
 462 in the NS or both. We have run permutations of all MSVs of the vertical geometry (Fig. 2) applied
 463 to the two geometries. The next two Sub-Sections show a summary of these hundreds of results
 464 into four different categories: 1) Those where both datasets push for the same model. We call this

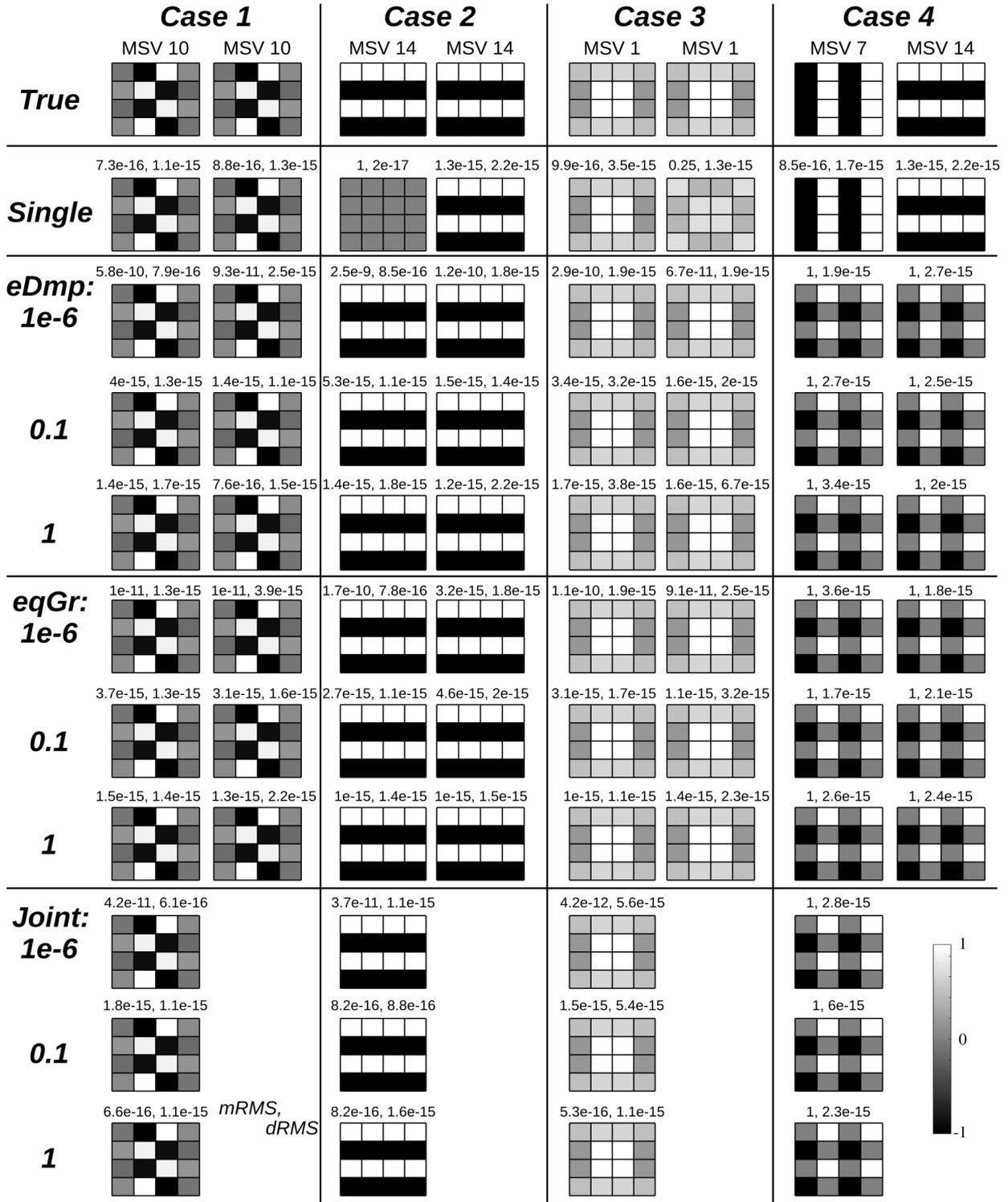
465 IS Corroboration. 2) Those where NS transfer occurs. 3) Those where IS Contradictions hinder
466 incorporation of MSVs. And 4) those where a mixture of the other three occurs simultaneously.

467 3.2.2 *MSV Inheritance Without Resistance*

468 Fig. 10 shows cases where MSVs are adopted without resistance, as they do not incur an increase
469 in the data misfit. Table 2 describes these results in detail (Cases 1 to 4). The best case scenarios
470 here all have the same true model for both geometries (Cases 1-3). MSV adoption without resis-
471 tance can come in two forms. The first has the two geometries corroborate the same result (IS
472 Corroboration; see case 1), i.e. the data for both geometries activate the same IS model compo-
473 nent. The other form occurs when Null-Space transfer fills in missing pieces (Cases 2 and 3). In
474 the worst case (4), mutual Null-Space Contamination causes the MSVs to combine into an incor-
475 rect model, without providing any signal in the data misfit. Note that all these forms of adoption
476 occur already at the lowest level of regularisation, because there is nothing in the respective data
477 that will resist it.

483 3.2.3 *Resisting Incorporation of MSVs*

484 Fig. 11 shows relevant cases (Case 5 to 7) where MSVs are resisted when they incur an increase in
485 the data misfit. Table 2 describes these results in detail. The incorporation of MSVs that produce
486 IS Contradictions is resisted in every form of coupling shown here. As regularization increases,
487 the contradicting MSVs are forced into the results. In consequence, the result tries to find a new
488 optimum which minimizes the data effects of this adoption. Because of this, the results start to
489 move away from both true models at high levels of regularisation. The Joint inversion of case 7
490 is especially interesting. Already at low regularization, the horizontal geometry is able to force a
491 model component through the NS into the final result. This is surprising, as the true model lies
492 completely in the IS of the vertical geometry. This shows, that resisting vectors due to IS Contra-
493 dictions may lead certain components to morph into the closest NS component. That this is not
494 always the case can be seen in Case 5. The behavior of the Joint inversion of Case 7 then contin-
495 ues quite similar to epoch damping and equivalent gradients. As regularization increases, the NS



478

479 **Figure 10.** Combined inversion cases 1-4 compared using single inversion, epoch damping, equivalent
 480 gradient constraints, and joint inversion. True model numbers correspond to Fig. 3. The two numbers above
 481 the estimated models correspond to the model and data RMS misfit (mRMS and dRMS), respectively. All
 482 estimated models for case 4 have values up to +/-2, except for the single inversions.

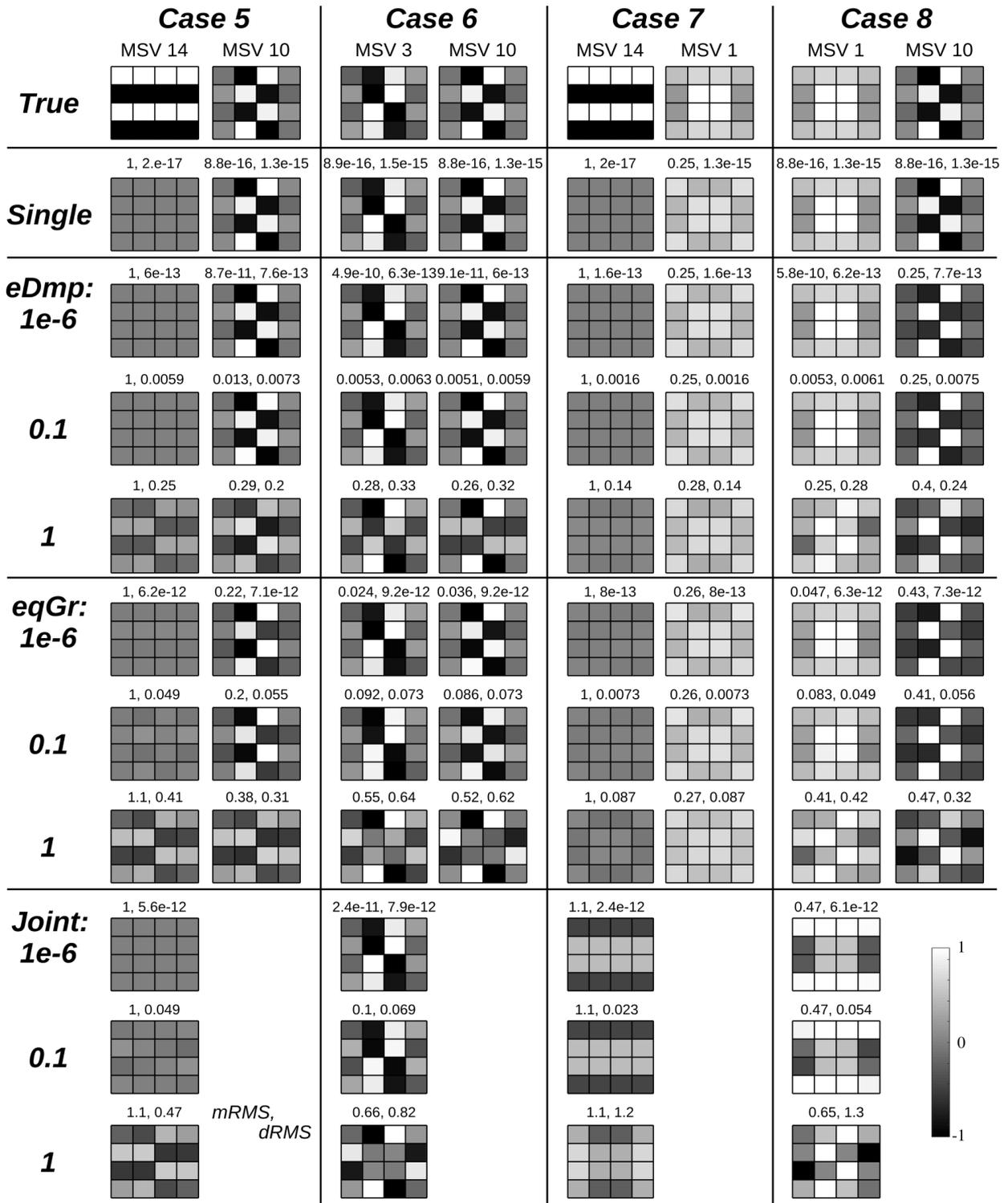
Table 2. Descriptions of the combined inversion behaviors observed in Figs. 10 and 11. **Geom:** Geometry, **MSV:** true Model-Space Vector, **VG:** in which of the Image Space (IS), Null Space (NS), or both (B) the true MSV lies for the vertical geometry, **HG:** in which of the IS, NS, or both the true MSV lies for the horizontal geometry

Case	Geom.	MSV	VG	HG	Behavior
1	V. uses	10	IS	IS	Both geometries corroborate the same result.
	H. uses	10	IS	IS	Neither affect the data misfit.
2	V. uses	14	NS	IS	The missing component for the first geometry is added without resistance. The 2nd geometry thereby completes the image.
	H. uses	14	NS	IS	
3	V. uses	1	IS	B	Similar to case 2 with part of the model being reproduced by the horizontal geometry already.
	H. uses	1	IS	B	
4	V. uses	7	IS	NS	Mutual null-space contamination. Both model vectors are summed together without impacting either data misfits.
	H. uses	14	NS	IS	
5	V. uses	14	NS	IS	The vertical geometry's inversion resists the inclusion of the IS model vector, as it contradicts its data. Increased regularization forces a mutual model, which does not correspond to either true model. The data misfit is strongly affected.
	H. uses	10	IS	IS	
6	V. uses	3	IS	IS	The inversions of both geometries resist the introduction of IS vectors which contradict their data. As in case 5, increased regularization leads to models which correspond to neither true models, and leads to increased data misfit.
	H. uses	10	IS	IS	
7	V. uses	14	NS	IS	The vertical geometry's inversion resists the inclusion of the IS model vector, as it contradicts its data. The combined inversion with the vertical geometry worsens the already poor results of the horizontal geometry.
	H. uses	1	IS	B	
8	V. uses	1	IS	B	The NS component of MSV 1 is adopted in the horizontal model without resistance (nor a signal in the data misfit). Resistance of contradictory IS components again leads to poorer results as regularization increases.
	H. uses	10	IS	IS	

496 component remains and other components are added due to the IS Contradictions. In Case 6 both
 497 geometries resist the IS components of the other geometry, requiring a high level of regularization
 498 before the models start deteriorating and resemble the Joint inversion model.

499 3.2.4 Combinations of Image-Space Contradictions, Corroboration, and Null-Space Transfer

500 As any true model can be obtained through linear combinations of MSVs, each of the three de-
 501 scribed behaviors can (and will for large problems) occur simultaneously. As an example, Case
 502 8 combines both NS transfer and IS contradictions. As seen with the other cases, the NS transfer
 503 occurs with very little regularization (see e.g. the mRMS for eDmp with $\alpha = 1e - 6$), whereas
 504 the IS contradictions cause the result to deform away from both true models with increased reg-
 505 ularization. Here again, the Joint inversion case struggles to deal with the contradictions in both
 506 datasets. For $\alpha = 1e - 6$, the result again adds the same NS component as in Case 7. This time it is
 507 combined with MSV 1. As regularization increases, MSV 10 and MSV 12 are major components
 508 here, with other components also coming in.



509

510 **Figure 11.** Combined inversion cases 5-8 compared using single inversion, epoch damping, equivalent
 511 gradient constraints, and joint inversion. True model numbers correspond to Fig. 3. The two numbers above
 512 the estimated models correspond to the model and data RMS misfit, respectively.

513 Geophysical inversions are generally not solved using a single MSV. Therefore, real investiga-
 514 tions using Joint, Constrained, and Time-Dependent inversion would always include some amount
 515 of Null-Space Transfer, Image Space Corroboration and Image Space Contradictions. In this sec-
 516 tion we have seen that NS Transfer occurs at real low levels of coupling between the two models.
 517 At the same time, IS Contradictions are harder to impose on the models, as they contradict the data.
 518 With the understanding gained in this Section, we may be able to identify each of these three cases
 519 by changing the regularization, and by comparing the results with those from single inversions.
 520 It is thus fundamental to perform individual inversions for comparison. Additionally, our results
 521 point out benefits of producing multiple models using constraints over jointly inverting for one, as
 522 well as additionally issues only encountered when jointly inverting for one model. We therefore
 523 suggest it is preferred to produce multiple models using a combined inversion method of choice
 524 (e.g. epoch damping, cross-gradient constraints, etc.), compared to Joint inversion. If practitioners
 525 choose to produce a single model using Joint inversion, however, we suggest it is vital to at least
 526 produce multiple models in a combined inversion of choice to check for the robustness of this
 527 jointly inverted single model. When looking at real problems, however, these issues are further
 528 complicated by data errors, which we will investigate next.

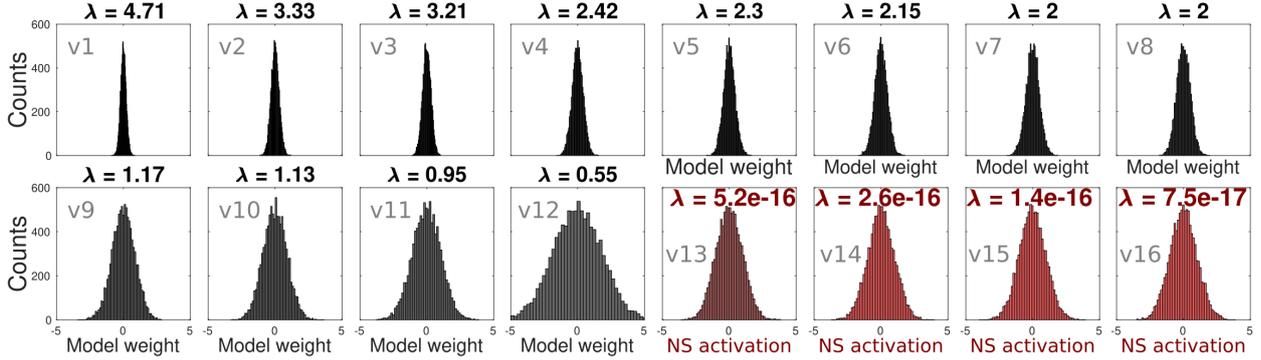
529 **4 PART II: THE INFLUENCE OF DATA ERRORS**

530 It is often repeated, that the influence of data errors scales with the inverse of the singular value.
 531 We put this into context by inverting random Gaussian errors and individual outliers. Based on the
 532 results, we describe how data errors influence Joint, Constrained, and Time-Dependent inversions.

533 **4.1 Repeated Investigation Using Gaussian Errors**

In this investigation, we used the vertical geometry (Fig. 2) to invert for random Gaussian errors
 10 000 times ($\mu = 0$ and $\sigma = 1$), without including any true data.

$$Gm = \vec{\sigma}_e \quad (8)$$



534
 535 **Figure 12.** Histograms of model weights for 10 000 iterations of random Gaussian errors. Each count
 536 corresponds to a single inversion. Each subplot corresponds to a model-space vector in Fig. 3, except for
 537 the last entries in red. Those depict the activation of the data-space vectors instead, as these vectors lie in
 538 the Null Space. λ denotes the associated singular value.

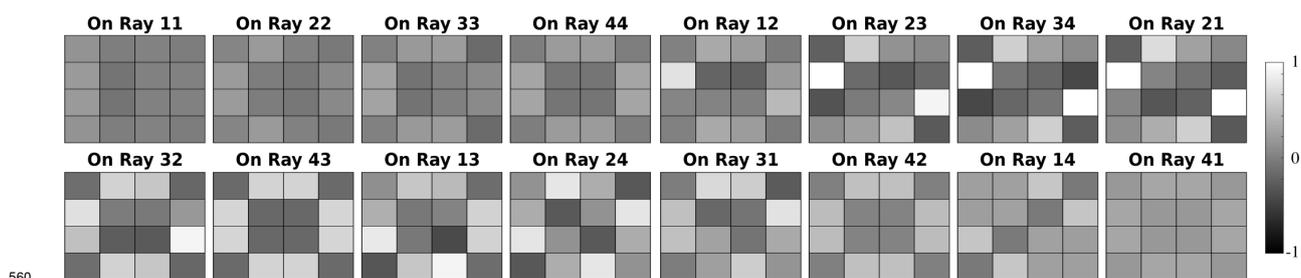
539 with, $\vec{\sigma}_e$, the random errors for each entry. Fig. 12 shows the corresponding model weights for
 540 each MSV.

541 Thus, repeated Gaussian errors produce a bell-curve distribution of these model weights with
 542 a mean around zero. As the singular value decreases, the width of the distribution (and thus the
 543 possible error) increases. In other words, the lower the singular value, the more likely it is that a
 544 set of errors will activate a given MSV, and the larger the activation. This also means that MSVs
 545 with lower singular values could see smaller activation in a single inversion, compared to MSVs
 546 with higher singular values. It is just statistically less likely to occur.

547 To dive a little deeper, random errors in the data will have components aligned with the data-
 548 space vectors. This includes being aligned with the vectors in the NS. Errors in the latter case
 549 would not affect the result, whereas errors aligned with DSVs in the IS are incorporated in the
 550 result with a corresponding decrease in data misfit. Using e.g. truncated SVD to reduce the number
 551 of MSVs is therefore a suitable way of reducing the influence of such errors, at the cost of not being
 552 able to incorporate those MSVs in the result.

553 **4.2 The Influence of Outliers**

The influence of outliers is commonly explained using a linear regression example. This exam-
 ple compares the fitting of a line to a number of points, with and without an outlier. When using
 least-squares regression, the outlier dominates this problem and the resulting line visibly shifts.
 Although this is an important example to illustrate the importance of managing outliers in gen-



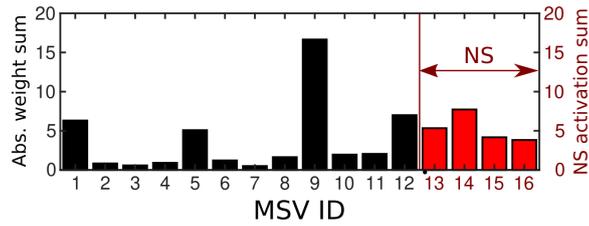
560 **Figure 13.** Model results due to outliers. Each subplot shows the final model obtained when only applying
 561 a 1 s traveltimes (no true data) on a single ray (e.g. R12, which connects shot 1 to receiver 2). The white cells
 562 in cases Ray 23, Ray 34, and Ray 21 shoot beyond the colorbar (max. 1.4).
 563

eral, this example is not representative of tomographic problems. Geophysical inversion as described here is rarely over-determined. (Changing the problem's discretization like in e.g. Multi-Dimensional Monte Carlo is one way of making a tomography problem over-determined; Sambridge & Mosegaard, 2002.) To better understand how outliers influence inversion results within the context of this paper, we have set up the following investigation: We apply a 1 s traveltimes (a single outlier) to one of the 16 rays in the vertical geometry and invert. Thus this investigation only looks at the direct influence of these outliers without further complicating the problem. Note that the outlier of 1 s is at the same level as the standard deviation of the Gaussian errors in the previous investigation. This is done to get normalized results and thus a comparison between the two investigations needs to be scaled according to a suitable size for an outlier. For investigation j , with the outlier on ray j , we get:

$$Gm = d_j^{outlier}, \quad \text{with } d_j^{outlier} = \begin{cases} 1 & \text{for } i = j \\ 0 & \text{for } i \neq j \end{cases} \quad (9)$$

554 Fig. 13 shows the 16 estimated models that this investigation produces. As none of the DSVs
 555 produce a single traveltimes, each outlier is reproduced by a linear combination of several DSVs
 556 (thus activating the corresponding MSVs). The linear combination of the activated MSVs show
 557 both destructive interference (e.g. rays 11 and 22), and undesired spikes (e.g. rays 23, 34, and 21).
 558 Also, MSVs 9 and 12 can clearly be recognized (e.g. in the results for rays 23, 34, and 21, MSV
 559 12 stands out, and an outlier on rays 32 and 43 clearly activates MSV 9).

564 Fig. 14 shows a summary of the corresponding model weights. Shown are the absolute sums of
 565 the weights for each MSV. (The individual results of this investigation are presented in Appendix



573

574 **Figure 14.** Summary of the outlier inversions. Each inversion has a 1 s outlier on one ray (no true data),
 575 which produces weights for each of the Model-Space vectors (MSVs). Shown are the sums of absolute
 576 model weights for each MSV ID. The red entries lie in the NS, and instead correspond to the sum of the
 577 corresponding Data-Space vector activations.

566 B.) The outliers preferentially activate MSVs 9, 12, 1, and 5, and there is quite some NS activation.

567 In the latter case, (part of) the outlier does not influence the estimated model. These results show,

568 that there is no relationship between the influence of the outliers and the singular value. For exam-

569 ple, the singular values of MSVs 9 and 10 are nearly the same (c.f. Fig. 3), but MSV 9 is obviously

570 much easier activated by an outlier. This further emphasises the importance of good handling of

571 these types of errors, as their effects on the final model is difficult to predict without specifically

572 testing for them.

578 4.3 How Data Errors Influence Joint, Constrained, and Time-Dependent Inversion

579 Section 3.2 showed how the activation of certain MSVs is hindered when they contradict the

580 data (IS Contradiction) and unopposed when they do not (NS Transfer). The exact same thing

581 happens with MSVs activated by errors. In the best-case scenario, model errors in one geometry

582 thus contradict the other geometry's data and are opposed, with a resulting increase in data misfit.

583 In the worst-case scenario, model errors are shared between geometries without resistance. Table 3

584 describes the influence of errors under different circumstances, including when true data is mixed

585 with data errors. For the cases where a single model is produced through Joint inversion, the

586 following needs to be taken into account on top of the entries in Table 3. Data errors that produce

587 IS Contradictions between each other, between themselves and the true data, or both can cause

588 additional NS entries to form (as seen in Case 7 in Fig. 11), as well as cause adjustments that far

589 exceed the influence of the errors on the single inversions (as seen in Case 8 in Fig. 11).

Table 3. Possible influences of errors in combined inversion. **EC(s)**: Error Component(s) **IS**: Image Space, **NS**: Null Space, **DSV(s)**: Data-Space Vector(s), **MSV(s)**: Model-Space Vector(s)

Given	When	Then
Common behaviors	1) ECs fall in the NS of the data	No influence on the models
	2) ECs produce IS Contradictions in the other geometry	MSVs adopted by own geometry and resisted by the other
	3) ECs corroborate the common MSVs	MSVs are adopted by both geometries without resistance
	4) ECs activate MSVs that lie in the NS of the other geometry	MSVs are adopted by both geometries without resistance
Gaussian errors in both datasets		Scenarios 1-4 apply (simultaneously)
One outlier for one geometry		Scenario 3 impossible Scenarios 1, 2, and 4 apply
One outlier each for both geometries		Scenarios 1-4 apply (simultaneously)
Errors and true data	ECs follow the 4 scenarios above	Similar behavior as above superimposed on the behaviors in Figs. 10 and 11
	ECs remove a true data component	True MSV removed from the result
	ECs combine with a true data component to activate a different DSV	True MSV replaced by a different one

590 5 DISCUSSION

591 Using visual descriptions and simple representative problems, we have incrementally built up a
592 set of fundamental behaviors of Joint, Constrained, and Time-Dependent inversion, with the goal
593 of providing an intuitive understanding of these abstract and complex methods. The core of these
594 fundamental behaviors can be reduced to two laws: 1) **Inherited model components that do not**
595 **affect the data misfit will be adopted without resistance.** 2) **Inherited model components that**
596 **affect the data misfit will be resisted.**

597 Behaviors like IS Corroboration and NS Transfer arise from the first law, whereas IS Contra-
598 dictions and component cleaning in non-linear problems arise from the second law. As we have
599 shown, these behaviors also apply to data errors, and the worst cases have data errors contaminate
600 one or both models without resistance. The first law also allows the usage of NS shuttles (e.g.,
601 Deal & Nolet 1996). Those, as we have discussed, are used to adjust existing models, whereas the
602 fundamental behaviors we have discussed are fundamentally part of the inversion process. Surpris-
603 ingly, Joint inversion sometimes produces NS components as a side product of the second law. We
604 suggest that this latter case cannot happen in non-linear inversion, as long as no additional con-
605 straints to e.g. an a-priori model are applied. Without any type of constraint, the inversion would

606 have no "power" to resist a change by an intermediate update to a model-component in the current
607 result. For Time-Dependent inversion, we have additionally discussed how changes in the geome-
608 try can lead to artificial changes in the results, even when the true model is the same. This is done
609 by having different MSVs in the IS and NS.

610 Translating these contributions to case-study-level problems will need further investigation.
611 Especially considering the massive 3D problems involving multiple geophysical properties, and
612 differences in parameterization, and differences in spatial/temporal resolution, and differences in
613 error levels and type, and the variety of workflow choices in pre-processing, initial models, and
614 during the non-linear iterations. Fully investigating model-vectors also becomes cost-prohibitive
615 very quickly as the size of the problem increases. Additionally, there are other avenues left unex-
616 plored within this range of simple representative problems. Nevertheless, having identified these
617 fundamental behaviors, we can now start applying them.

618 The rest of this discussion will therefore focus on the following: First, we make hypotheses
619 of how the presented behaviors would interact when using the presented methods simultaneously
620 (e.g. combined non-linear inversion with data errors). Based on that discussion, we will revisit the
621 benefits of producing multiple models simultaneously, over producing a single model with Joint
622 inversion. Then we will dive a little deeper into the relevance to other geophysical methods. Lastly,
623 we will show how inversion results can be improved knowing these fundamental behaviors.

624 **5.1 Predicting Behaviors of Combinations of the Presented Methods**

625 It is beyond the scope of this paper to fully investigate the interactions between the presented phe-
626 nomena. Though we may not be able to provide a complete picture without such an investigation,
627 the fundamental nature of the presented behaviors does allow us to make some predictions about
628 these interactions below. The ability to setup such hypotheses based on these fundamental behav-
629 iors is one of the main contributions of this work. The most important point for the following cases
630 is that changing the problem (by regularization, changing the geometry, or otherwise) will change
631 what data components map onto the IS and NS.

5.1.1 *The Influence of Data Errors in Smoothed Inversion*

This form of regularization extends the data space (Fig. 7) leading to increased opportunities for data errors to map onto the NS. Therefore, even though the regularized problem can add MSVs that are not contained in the true model (Fig. 6), it should make the inversion more robust against certain data errors.

5.1.2 *Combined Inversion Using Smoothing Constraints without Data Errors*

Smoothing regularization forces additional model components into each geometry's result (through vector re-orientation). Which components are added will differ between the two geometries. Some of these added components will thus transfer through the NS. These additional differences will also lead to increased opportunity for IS Contradictions. The model components that cause the added IS contradictions are coupled to the true model components. Resisting the contradictions will thus mean resisting the true model components. The strength of this resistance will depend on the size of the data contradictions, and on the level of smoothing. The exception here would be Joint inversion. The smoothing regularization would be added below the two G-kernels in Eq. 6. Thus, the components introduced by smoothing would not lead to an introduction of a corresponding component in the other geometry.

5.1.3 *Combined Inversion Using Smoothing Constraints with Data Errors Included*

Here, again, the smoothing constraints increase the likelihood of data errors mapping onto the NS (also in Joint inversion). For multi-model producing methods, the increased possibility for IS Contradictions (due to vector re-orientation) should also increase the likelihood that the inversion resists the inclusion of model components due to data errors.

5.1.4 *The Influence of Data Errors in Non-Linear Inversion*

As non-linear inversion changes what data components map onto the IS and NS when changing the geometry, we would expect that data errors may swap between mapping onto the NS and onto the IS. Thus, data errors could activate MSVs in the intermediate results that transfer through the

657 NS into the final models. Stated differently, data errors that have already activated components in
658 the intermediate model may no longer show up in the residual for the inversion to act upon. The
659 main issue here, is that data errors could thus influence the path that the inversion takes to the final
660 result. On the other hand, data errors that map onto the NS for a given geometry, are likely still
661 in the residual for the next geometry. This will always depend on the specific case, as data errors
662 that map onto the NS in one intermediate geometry, may already be satisfied by the existing model
663 in the next geometry. If data errors remain in the residual, those that map onto a geometry's IS
664 could remove (part of) MSVs produced in the previous result(s) that might otherwise have been
665 transferred through the NS. Data errors could thus also have a cleaning function. The likelihood of
666 this occurring depends on the size of the data errors, compared to the data components produced
667 by such a MSV and thus relates directly to that MSV's singular value (smaller singular value =
668 more likely).

669 *5.1.5 Smoothed Non-Linear Inversion Without Data Errors*

670 As the geometry changes, the MSVs required to fulfil the smoothing constraints will also change.
671 Thus, these constraints could produce additional components that lie in the NS of the subsequent
672 geometries. At the same time, the newly required components could contradict with components
673 in the previous model and thus (partially) remove them. Thus, smoothing could influence the path
674 towards the final solution in both positive and negative ways.

675 *5.1.6 Smoothed Non-Linear Inversion With Data Errors*

676 Smoothing increases the likelihood that certain data-error components are sent to the NS. As these
677 components likely remain in the residual for the next iteration, these components would repeatedly
678 have the opportunity to affect the IS of the result, except if they are always sent to the NS. Thus,
679 smoothing could lessen the influence of errors, at the cost of the added model components required
680 to fulfil the constraints. At the same time, the effects of the data errors and smoother could combine
681 and push the inversion path into an undesired direction.

682 *5.1.7 Combined Non-Linear Inversion without Data Errors*

683 In combined non-linear inversion, NS components can now also come through the coupling, and
684 in the case of Joint inversion, be produced due to IS Contradictions. Similar to non-linear inversion
685 of a single problem, Joint inversion can (partially) remove such components when contradictions
686 are encountered in either dataset. In contrast, in problems with combined inversion methods that
687 produce multiple models, each sub-problem can only clean their own components. Most problem-
688 atically, the coupling could allow a component to be reintroduced into a model after a previous
689 cleanup. To explain, we will discuss the case of two non-linear problems, A and B, that are coupled
690 using a combined method that produces a model for each sub-problem. These two sub-problems
691 can swap NS components as follows: 1) a component in result A transfers to result B during one
692 iteration, 2) in a subsequent iteration, problem A cleans out this component as it contradicts the
693 data, whereas it happens to remain in the NS of result B, 3) in a later iteration, problem A once
694 again orients this component in the NS, thus 4) allowing this component to be reintroduced into
695 problem A from problem B.

696 *5.1.8 Combined Non-Linear Inversion with Data Errors Included*

697 Here again, the inclusion of errors increases the possibility of IS Contradictions. Resisting the
698 corresponding model components pushes the associate data-error components into the NS. There,
699 they would be available for the next iteration to produce IS Contradictions again. Also for this
700 case, data-errors could influence the path of the non-linear inversions. This is especially likely
701 when data-errors produce IS Corroboration.

702 *5.1.9 The Last Two Combinations and Other Cases*

703 We invite the reader to hypothesize on how the last two combinations would behave. These are: 1)
704 Combined Non-Linear Inversion Using Smoothing Constraints without Data Errors, and 2) Com-
705 bined Non-Linear Inversion Using Smoothing Constraints with Data Errors Included. We also in-
706 vite readers to hypothesize on the following cases, that go beyond the presented investigations: 3)
707 Combined inversions joining three datasets into a single estimated model (e.g. electromagnetics).

708 4) Combined inversions that produce three or more estimated models (e.g. V_p , V_s , and density
709 in full-waveform inversion, or epochs in time-dependent inversion). 5) Pairwise inter-model con-
710 straints in time-dependent inversion, i.e., m_1 with m_2 , m_2 with m_3 , etc.

711 **5.2 Benefits of Producing Multiple Models Simultaneously**

712 The above interactions once again point out the benefit of producing multiple models with a form
713 of constraint between them. This firstly allows us to differentiate between NS Transfer and IS
714 Contradictions through the simple act of comparing models at different levels of regularization.
715 NS Transfer appears at very low level of regularization, whereas IS Contradictions increases with
716 increasing levels of regularization. Then, it helps decrease the influence of data errors on the final
717 results. Lastly, it allows us to check the robustness of features using data subsets, which is always
718 the case for Time-Dependent inversion (see e.g., Hobé et al. 2021).

719 **5.3 Relevance to Other Geophysical Methods**

720 Irrespective of parameterization type (e.g. cell, nodal), if there are the same amount of parameters
721 for the model, then every possible model can be deconstructed with the same MSVs. These MSVs
722 would be distributed differently when the physics changes, similar to how a change in geometry
723 redistributes the MSVs among the IS and NS. Thus, the insights from this paper can be applied
724 directly to investigations where the same method is used with different datasets.

725 A non-seismic example where the insights of this paper apply is Joint, or Constrained inversion
726 of two different arrays in ERT (e.g. Horo et al. 2021). The physics of ERT would lead to a different
727 redistribution of MSVs among the IS and NS, compared to the seismic example. Similarly to the
728 presented investigations, the problem corresponding to the first ERT array would have MSVs in
729 the NS, that lie in the IS of the second problem, and vice versa.

730 A translation of these fundamental behaviors is required, however, when coupling geophysical
731 methods using an empirical relationship (e.g. gravimetry and seismics; Haber & Holtzman-Gazit
732 2013 and references therein) and/or constraints between models with different parameterizations.
733 Filtering behaviors like IS Contradictions, NS Transfer, and IS Corroboration through an empiri-

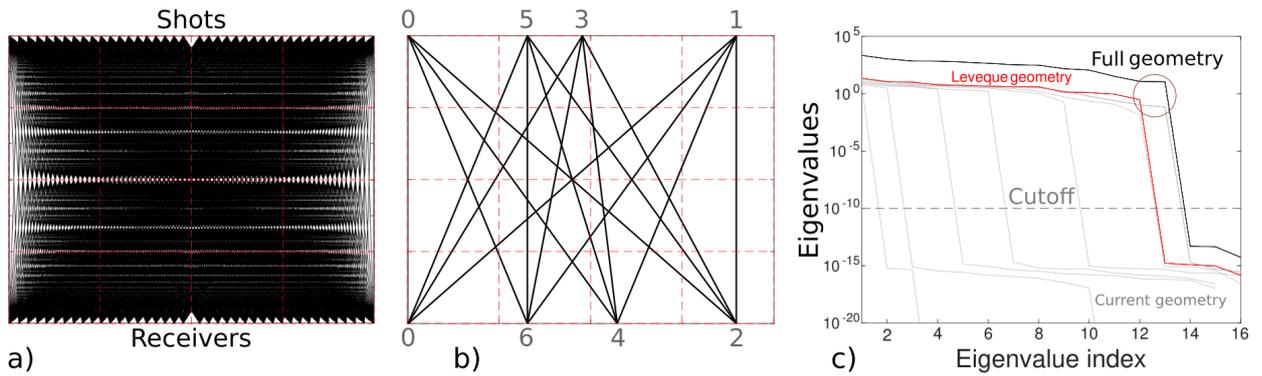
734 cal relationship will make it much more difficult to identify which part of one sub-problem's model
735 influences which part of the other model. However, producing multiple models using different lev-
736 els of regularization, should also allow us to differentiate between these behaviors. Additional
737 issues may arise when the two sub-problems have different parameterizations. Without further
738 investigation, it is not clear to us how coupling two problems with a different amount of MSVs
739 would affect results. Does a MSV of the sub-problem with less parameters affect multiple MSVs in
740 the other sub-problem? If yes, what happens when both IS and NS components are activated at the
741 same time? If not, are there components in the sub-problem with more parameters that are never
742 affected by the other sub-problem, and do those, *vis versa* never affect the first sub-problem?
743 Additional difficulties would arise trying to predict these behaviors for methods that change their
744 parameterization during the non-linear inversion process.

745 **5.4 Improving Inversion Results**

746 In this Section, we will describe multiple possibilities for alleviating the severity of Null-Space
747 Contamination and the influence of errors described above. We subdivide these options in the
748 following categories, which we will discuss below: 1) Optimized Experimental Design. 2) Using
749 more and repeater data. 3) Preventing Null-Space Contamination. The most obvious fourth cat-
750 egory is the identification, reduction and removal of errors. For this final category, we refer the
751 reader to the literature associated with the individual geophysical methods. One example that re-
752 duces picking errors before picking even begins is the "shift and stack" method in active seismics
753 (Park et al. 1996).

754 *5.4.1 Optimized Experimental Design*

755 Lévêque et al. (1993) used the Leveque geometry to great effect to explain the specifics of checker-
756 board tests. This geometry has also been very helpful in this work to visualize the fundamental
757 behaviors we are interested in. However, this geometry is not the optimal geometry for the 4x4
758 cell models in question. Here, we will show how all the results in this paper could be improved



768

769 **Figure 15.** Optimized-Experimental-Design results. **a)** Geometry of all sampled rays. **b)** Optimal geometry
 770 obtained for this setup with rays numbered in order of addition (the two zero positions were defined in
 771 advance). Red lines in a) and b) denote cell boundaries. **c)** Eigenvalue spectra comparison. The eigenvalues
 772 (γ) relate to the singular values (λ) as: $\gamma_i = \lambda_i^2$. The spectra of each OED iteration are shown in grey. The
 773 spectrum of the Leveque geometry is shown in red, whereas the spectrum using all rays in a) is shown in
 774 black. Eigenvalues below the cutoff are seen as being in the Null Space. The dark-red circle emphasizes the
 775 differences of the three spectra for the lowest eigenvalue in the Image Space.

759 by optimizing the placement of the stations and receivers using Optimized Experimental Design
 760 (OED; e.g., Maurer et al. 2010).

761 At the start of our OED study, we define all possible instrument placements and produce the
 762 corresponding rays (Fig. 15a). After defining a starting point (to reduce the computational cost)
 763 with a first shot and a first receiver (0 and 0 in Fig. 15b), we search through the remaining shot
 764 locations, to find the shot which will maximize the normalized eigenvalue spectrum (Fig. 15c) of
 765 the current step (though other criteria exist; Curtis 1999a,b; Routh et al. 2005; Ajo-Franklin 2009;
 766 Maurer et al. 2010). This process is repeated, while alternating between shots and receivers, till
 767 the desired number of shots and receivers are obtained.

776 When comparing the three eigenvalue spectra (Fig. 15c) of the Leveque geometry, of the full
 777 geometry (using all rays), and the "optimal" geometry obtained in this way (Fig. 15b), we observe
 778 the following: 1) The optimal geometry using 4 shots and 4 receivers elevates one additional
 779 singular value above the cutoff compared to the Leveque geometry. 2) The optimal geometry has
 780 the same number of MSVs in the IS as the full geometry. 3) The addition of additional rays beyond
 781 the optimal geometry further increases the absolute eigenvalues. 4) The number of MSVs in the
 782 Image Space obtained using the Leveque geometry can also be obtained using less receivers. Note
 783 that the normalized eigenvalue spectrum of the optimal geometry and the full geometry are almost
 784 equal (not shown).

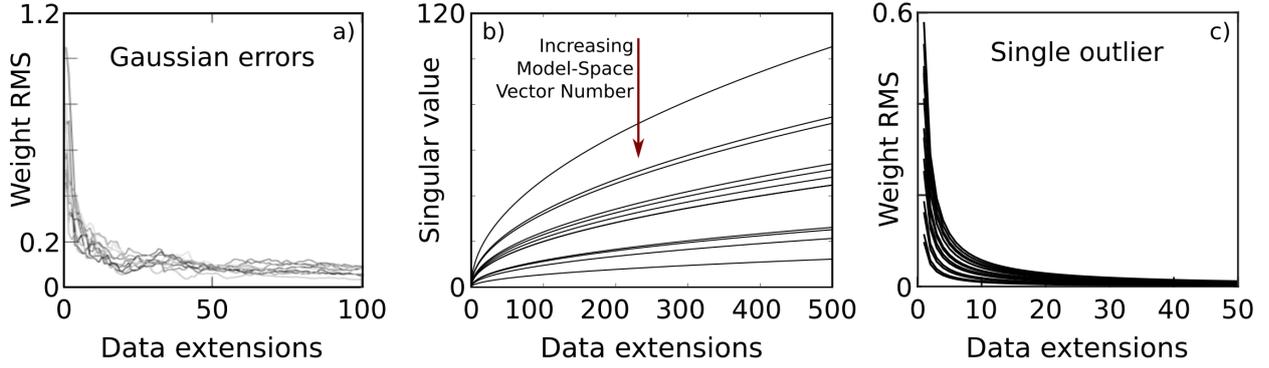
785 To connect back to the influence of errors, OED directly reduces the influence of errors by im-
 786 proving the eigenvalue spectrum. This effect is best seen in Fig. 12, where increasing the singular
 787 value leads to a narrowing of these bell curves. Additionally, the elevation of additional MSVs into
 788 the IS reduces the opportunity for NS Contamination, and increases the possibility of IS Contra-
 789 dictions. Lastly, the elevation of the entire eigenvalue spectrum for the full geometry shows that
 790 the optimal geometry is only optimal from the point of view of the normalized eigenvalue spec-
 791 trum. The higher spectrum of the full geometry, compared to the optimal geometry, corresponds to
 792 a smaller standard deviation of model weights due to Gaussian errors. Similar improvements are
 793 expected for the other optimization measures in OED, with these improvements being more local-
 794 ized for measures meant to improve specific model features (e.g. Routh et al. 2005; Ajo-Franklin
 795 2009). Because adding instruments is often more expensive, compared to adding more shots, we
 796 will now look at the benefit of adding shots on the same locations for the same instruments.

797 5.4.2 *The Benefit of Additional Data*

798 To show the benefit of additional data, we extend the investigations into the influence of errors
 799 as follows. We repeat the rays of the Leveque geometry for each data extension. Then, we either
 800 apply Gaussian noise to each ray, or a single outlier on one of the first 16 rays. This is formalized
 801 in Eqs 10 and 11 (for investigation j):

$$\begin{bmatrix} G \\ G \\ \vdots \\ G \end{bmatrix} m = \vec{\sigma}_e \quad (10)$$

$$\begin{bmatrix} G \\ G \\ \vdots \\ G \end{bmatrix} m = d_j^{outlier}, \text{ with } d_j^{outlier} = \begin{cases} 1 \text{ for } i = j \\ 0 \text{ for } i \neq j \end{cases}, \text{ for } j < 17 \quad (11)$$



805

806 **Figure 16.** Results for the error investigations when extending data. Each data extension adds the same 16
 807 rays. In **a)** random Gaussian errors are applied to the available rays as traveltimes (errors not repeated). The
 808 different curves correspond to ten different seeds. **b)** Shows the Singular values as data increases. In **c)**
 809 each curve corresponds to a single outlier being applied to one of the first 16 rays. All other rays for each curve
 810 have a 0 s travelttime.

Fig. 16 shows the results. Surprisingly, the influence of errors behaves like an L-curve when using such data extensions. For the first few extensions, the impact on the estimated model rapidly decreases for both investigations. Although the model misfit continues to decrease as more data is added, the rate of decrease declines rapidly. For the outlier investigation, this influence can be summarized to the following equation:

$$R(n) = R_{init}/n \quad (12)$$

802

Here, R , is the RMS of all the model weights, R_{init} , is the initial value of R , corresponding to the Leveque geometry, and n , is the number of data extensions ($n = 1$ corresponds to the Leveque geometry).

804

Fig. 16b shows how the individual singular values improve as we repeat data on the same rays. This result corresponds to both the Gaussian error and outlier investigations. Here, the ability to increase the singular value by adding data on the same rays is clearly correlated with the initial value. A line fit provides the following relationship (Eq 13):

$$\lambda(n) = \lambda_{init} * \sqrt{n} \quad (13)$$

811

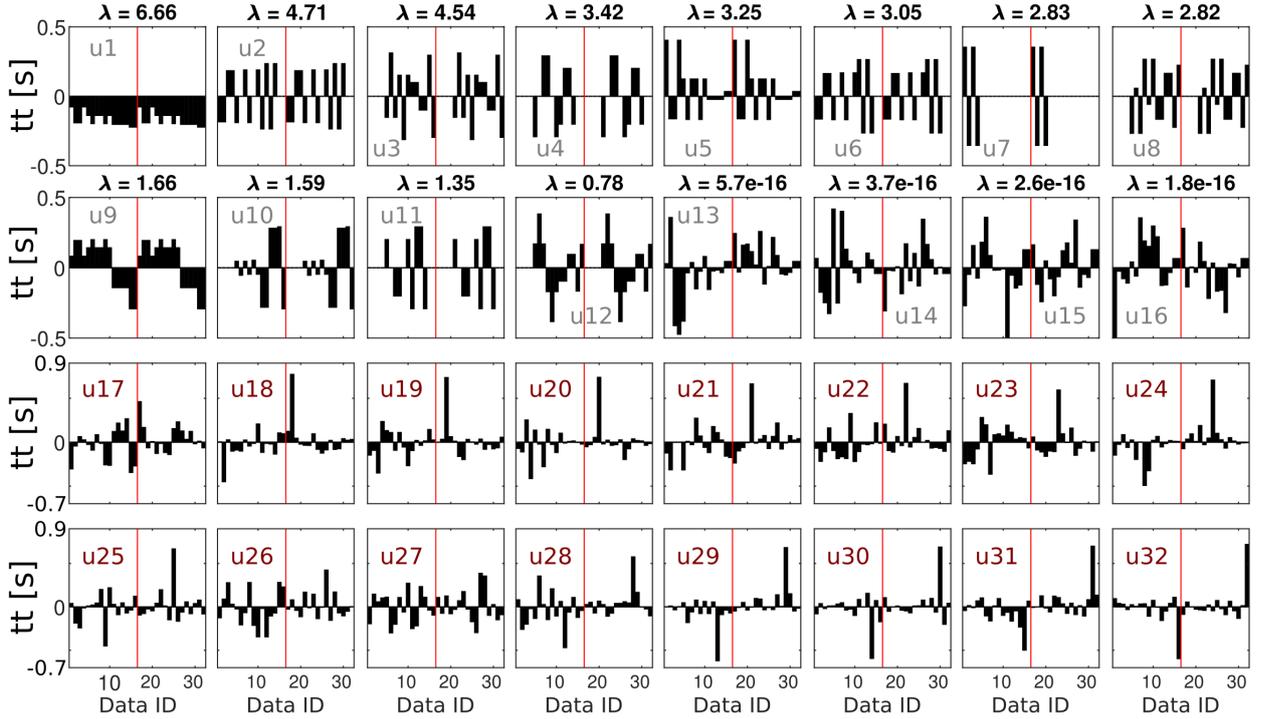
Although the largest initial singular value grows much larger compared to the others, this value was already quite robust against errors to begin with (Fig. 12). The smallest initial singular value (lowest curve) increases rapidly at the beginning, after which the rate of increase declines. As

813

814 the lower-most curves are the most sensitive to errors, their trend may be enough to explain the
815 L-curve behavior in Fig. 16a and c: While these lower-most curves grow rapidly, the influence of
816 the errors on the most sensitive singular values rapidly decreases. As the growth of these curves
817 decreases, the reduction in error sensitivity decreases as well. Additionally, Eq. 13 means that
818 the conditioning number (largest Singular value divided by the smallest Singular value above the
819 cut-off) remains equal when using extensions with the same rays. Whether this is applicable for
820 extensions using specific subsets of rays will require further investigation.

821 We show a different aspect of the reduction in error sensitivity in Fig. 17. It shows the DSVs
822 of the case where the rays are sampled twice. While there are 16 MSVs, with four in the NS, there
823 are 32 orthogonal DSVs (one for each traveltime). Each of these vectors has 16 new entries, which
824 are a duplicate of the first 16 for the DSVs in the IS (u_1 - u_{12}). Most of u_{17} - u_{32} instead have a
825 single large spike on the right of the red line, along with smaller entries. The relative magnitude
826 of this single spike compared to the other entries increases as we include more data. As there are
827 no singular values associated with these latter 16 DSVs, all of these vectors map onto the NS, and
828 thus do not affect the model. Additional data thus reduces the sensitivity to errors, because there is
829 a greater likelihood for them to map onto the NS (a linear combination of the DSVs can reproduce
830 any data vector). As outliers are unlikely to show up for every repeat observation, it is highly likely
831 that they will map onto one of the large spikes of the DSVs that map onto the NS. Whether there
832 is a maximum amount of outliers per data amount, or whether systematic outliers are more likely
833 to map onto the IS will require further investigation.

838 As we have shown here, additional data massively reduces the influence of both random Gaus-
839 sian errors and of single outliers. This reduction occurs through the combined effect of increasing
840 the singular values, and by capturing discrepancies to the IS vectors in an extended NS. The same
841 general behavior is expected from additional data, which does not repeat a previously sampled
842 geometry (e.g. the other rays in Fig. 15a). As Figs. 16a and c look like an L-curve (e.g., Menke
843 2018), we suggest that such trade-off curves can be used to identify both the minimum and opti-
844 mum amount of data to collect, as part of an experimental design study. Special versions of OED
845 exist that allow users to optimize for multiple things for a given budget (e.g., Maurer et al. 2010).



834

835 **Figure 17.** Data-Space vectors for the case where the geometry is duplicated. The red line indicates the start of
 836 the duplicated values. Vectors 13-32 map onto the Null Space as there is no singular value (λ) connecting
 837 them to a model-space vector.

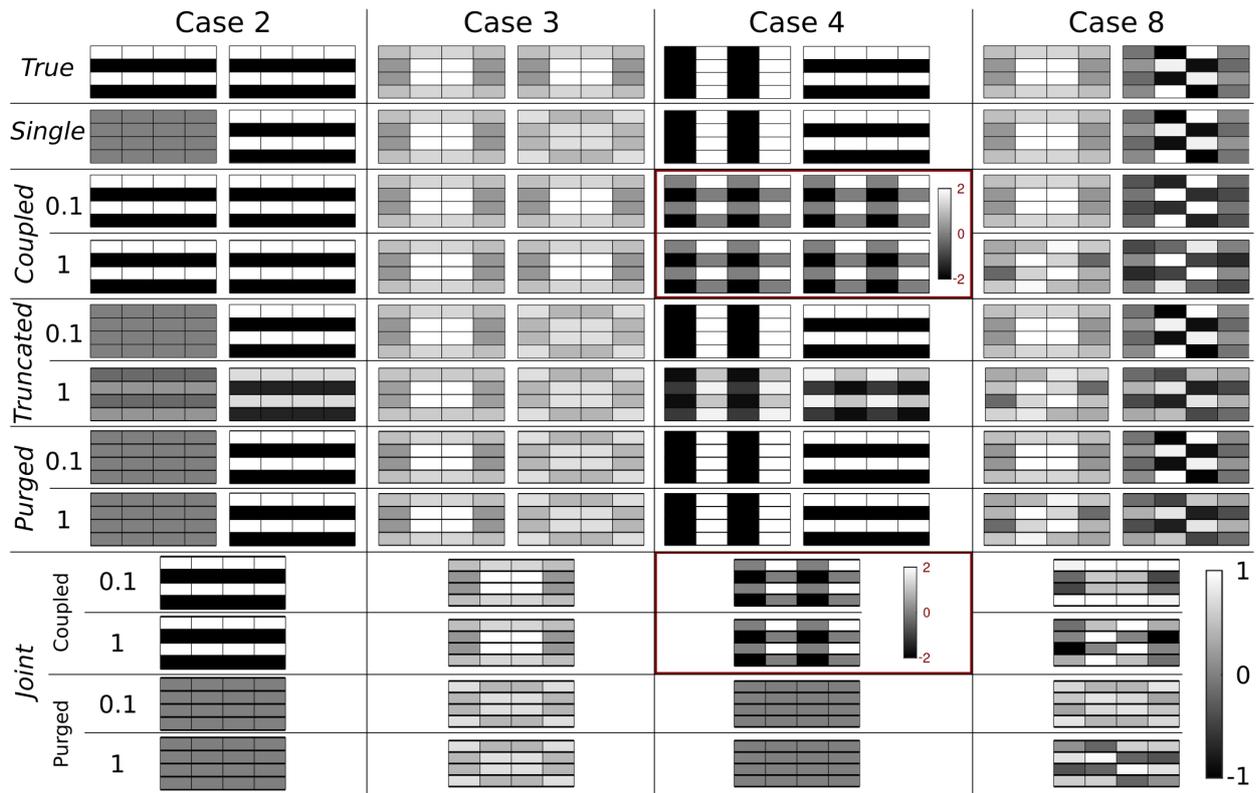
846 Thus one could create a method that finds the optimum amount of data for the optimum amount of
 847 instruments given the budget and prior knowledge of the area under investigation. To our knowl-
 848 edge, producing an L-curve to identify the optimum amount of data for a given geometry has not
 849 previously been discussed.

850 5.4.3 Preventing Null-Space Contamination

851 NS Transfer is beneficial when it fills out missing pieces of the true model (given equal true
 852 models). However, it also allows the very worst cases of the combined inversion methods (e.g.
 853 Case 4 in Figs. 10, and table 2). These cases of NS Contamination become possible, because
 854 combined inversion couples these unwanted NS components in one geometry, with IS components
 855 in the other geometry. Thus, when the IS components are activated by data in one geometry,
 856 the coupling introduces these NS components into the other geometry. NS Contamination can
 857 therefore be prevented completely, by removing the coupling of NS components before calculating
 858 the estimated models. We discuss three such decoupling strategies in Appendix C, along with a
 859 deeper look into how this coupling takes place. We will summarize these strategies here, and show

860 some initial results (Fig. 18). All three strategies start with a comparison of the MSVs in the
861 combined inversion with those in the single inversions for both geometries. The first method uses
862 this comparison to do something similar to truncated SVD, as we truncate coupled MSVs who's
863 entries lie in the NS for both geometries. An additional cleaning step is used when one geometry
864 has more entries in the IS, compared to the other geometry (see Appendix C). In the second
865 method, we use the comparison to purge the NS components of the single inversions from the
866 coupled MSVs. The third method produces the same result in the opposite way. Here, we extract
867 the components that lie in the IS of the single inversions from the coupled MSVs. Fig. 18 compares
868 results of these three strategies against the results of methods where the coupling is intact. Fig. 18
869 does not show the truncation strategy for Joint inversion. The additional complications arising for
870 this method make this strategy less suitable compared to the purging and IS retention strategies.
871 These latter strategies also need different implementations for Joint inversion because of these
872 complications (see Appendix C).

882 All three strategies do well preventing NS Transfer for $\alpha = 0.1$. (All results of the NS purging
883 strategy are equivalent to those IS retention method.) Although the truncation strategy is the most
884 practical, Fig. 18 shows that it does not clean out NS Contamination entirely for large α values.
885 The poorer results of the truncation strategy are especially clear for Cases 2-4 ($\alpha = 1$). Case 8
886 shows the benefit of proper decoupling when trying to differentiate between NS Transfer and IS
887 Contradictions. The coupled results using $\alpha = 1$ include the NS components, whereas the purged
888 results using $\alpha = 1$ only include components in the IS of both geometries. Case 4 shows the added
889 complexity when decoupling of Joint inversion. As Joint inversion only produces one model, we
890 need to choose which NS components to decouple. This can be only those of one geometry (two
891 choices), and the NS components of both geometries at the same time. These choices would also be
892 possible for methods that produce multiple models. In Joint inversion, however, the consequences
893 are more severe. In Fig. 18 we present the Joint inversion strategy, that purges the NS components
894 of both geometries. In Case 4, this results in both true models not mapping onto the final result.
895 For Case 3, only the component in the IS of both geometries is retained. Case 8 shows, that also



873

874 **Figure 18.** Inversion results comparing decoupling strategies with commonly used combined inversion
 875 methods. The top two rows show the True models (one each for the vertical and horizontal geometries), and
 876 the results for single inversions of these true models. Then, *Coupled*, *Truncated*, and *Purged* correspond to
 877 those strategies applied to inter-model minimization (e.g. epoch damping). The adjacent numbers are the
 878 applied regularization factors ($\alpha = 0.1$ and $\alpha = 1$). The bottom two rows show the Joint inversion cases.
 879 The strategy that retains the IS components produces equivalent results to the purging strategy. Note that
 880 the colorbars for the results in the red rectangles differ from all other results, to emphasize the worst case
 881 scenario.

896 the NS components produced due to IS Contradictions are not retained in the final result (also see
 897 Case 7 in Fig. 11).

898 There is one major downside to this decoupling strategy. As it negates NS Transfer between
 899 geometries in combined inversion, decoupling not only stops NS Contamination (Case 4 and 8 in
 900 Fig. 18); It also stops the transfer of missing components that are common between the true models
 901 (Case 2 and 3 in Fig. 18). Nevertheless, the ability to produce separate results with and without
 902 NS transfer is very promising. Although it is beyond the scope of this paper, these strategies will
 903 need more investigation, even beyond what is shown in C. We suggest that the IS retention strategy
 904 has the largest potential. The biggest benefit of decoupling lies in its ability to prevent data-errors
 905 in one dataset to transfer through the NS to another sub-problem without resistance. Additional
 906 benefits arise in Time-Dependent inversion, where decoupling allows us to focus on differences

907 supported by the data. Lastly, the results in Fig. 18 further emphasize the benefits of producing
908 multiple models using some form of regularization between them, over a single model using Joint
909 inversion.

910 **6 CONCLUSIONS**

911 Using a visual analysis of simple inverse problems, we have shown fundamental behaviors of
912 common inverse methods. Although these problems use traveltimes, the results are applicable
913 to all geophysical methods that use inversion. Before tackling the complex cases, we have, e.g.,
914 described the fundamental workings behind the commonly used L-curve, and uncovered the unde-
915 sired creation of Null-Space (NS) components within non-linear inversion. For Joint, Constrained,
916 and Time-Dependent inversion, we have shown the following three fundamental behaviors: Image-
917 Space (IS) Corroboration, IS Contradictions, and NS Transfer. These fundamental behaviors lead
918 to the following best and worst case scenarios under specific conditions. Best Cases: 1) The recon-
919 struction of equal true models is improved, 2) mixing of unequal true models is resisted, and 3)
920 errors map onto the NS. Worst Cases: 1) Errors map onto the estimated model(s), 2) errors transfer
921 through the NS without resistance, 3) unequal true models mix without affecting the data mis-
922 fit, and 4) undesired NS components are created when IS contradictions occur in a single model
923 producing Joint inversion.

924 Based on these insights, we propose multiple avenues to improve inversion results. Next to
925 the common practices for error and outlier removal/reduction, we show that inversion results can
926 be improved using Optimized Experimental Design (OED), incorporating more data, and by de-
927 coupling the NS components in the Joint, Constrained, and Time-Dependent inversions. In this
928 context, we present a novel addition to OED, whereby the optimum amount of data can be found
929 after finding the optimum instrument placements, through the use of a trade-off curve showing the
930 influence of errors against the number of data extensions. The decoupling methods are especially
931 useful for Time-Dependent inversions, as it allows us to focus on differences supported by the
932 data, without the influence of NS Transfer.

933 Overall, we show the benefits of producing multiple estimated models using constraints over

934 combining the inverse problems into a single model. The insights from this paper have the poten-
935 tial to fundamentally improve multiple aspects of geophysical survey design, data pre-processing,
936 model interpretation, and the creation of new inverse methods. More importantly, the visual analy-
937 sis of simple inverse problems should greatly reduce the difficulty for newcomers and practitioners
938 to improve their intuition of these complex and abstract methods.

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944 to provide his student's with an intuitive sense of geophysical inversion lead to seeing the massive
945 contribution in Lévêque et al. (1993), which further inspired the investigations in this manuscript
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949 **8 DATA AVAILABILITY STATEMENT**

950 No data nor code is made available for this manuscript. We highly recommend reproducing these
951 simple investigations from scratch.

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1058 **APPENDIX A: IN-DEPTH SMOOTHING CONSTRAINTS INVESTIGATION RESULTS**

1059 In Fig. 8 (Section 3.1.4), we only showed the model weights for four reconstructions using smoothed
1060 inversion. The results for all 12 data-producing MSV are shown in Figs. A1 and A2.

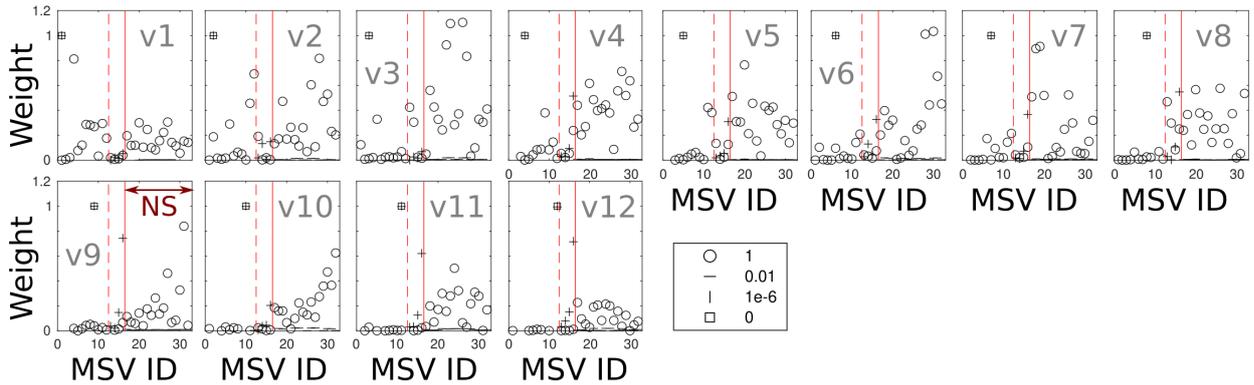


Figure A1. Inversion results for reconstructing true models (V1-V12 here) in Fig. 3 using the vertical geometry and different levels of smoothing. V13-16 are not shown as they do not produce traveltimes. To the left of the full red line, these subplots show the absolute weights for each of the model-space vectors (MSVs) of the smoothed vertical problem, i.e., these values show how much of each MSV (values on the x-axes) is used when inverting for a given true model (e.g. V1) using different levels of smoothing. The MSV IDs to the right of the full red line instead show how much of the data is captured by the Null Space (NS) of the data space, i.e. the absolute data length in the direction of that DSV in the NS. Values between the broken red line and the full red line correspond to the four MSVs that are in the NS for the original problem. These are no longer in the NS for the smoothed problem, as small Image-Space components have been added that now activate these MSVs using data. Here, all values have been truncated below $1e-3$.

1061 **APPENDIX B: IN-DEPTH OUTLIER INVESTIGATION RESULTS**

1062 For story purposes, we did not go into depth in Section 4.2. Here, we present and discuss the
 1063 individual results of the single outlier investigation using only the vertical geometry. As a reminder,
 1064 we apply a 1 s traveltime to a single ray (a single outlier) and invert for that data without adding
 1065 any other data.

1066 Fig. B1 shows the resulting weights on the individual MSVs for each affected ray. Here, the
 1067 rays are ordered as implemented from left to right, with the two numbers corresponding to the shot
 1068 and receiver that the ray links together (e.g. R34 has shot 3 and receiver 4). As none of the DSVs

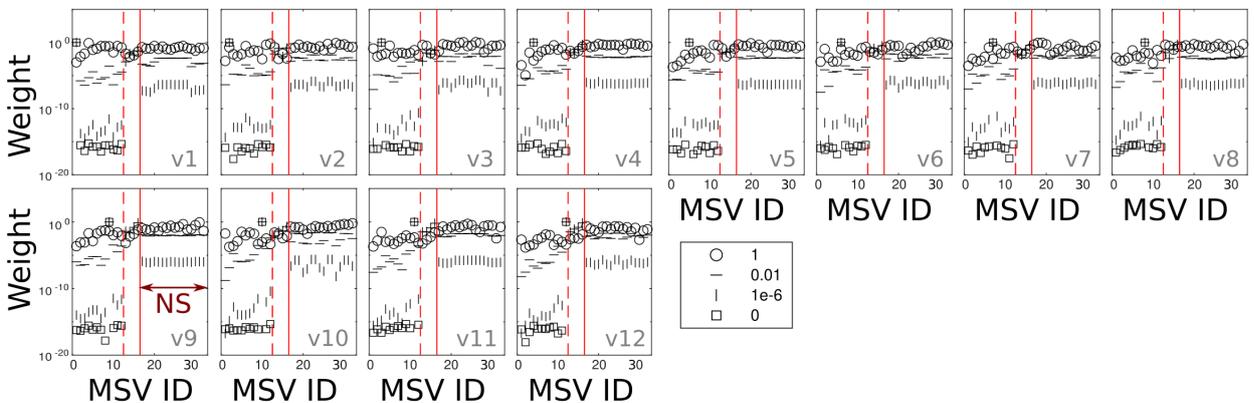


Figure A2. Log-linear view of Fig. 8 without truncation.

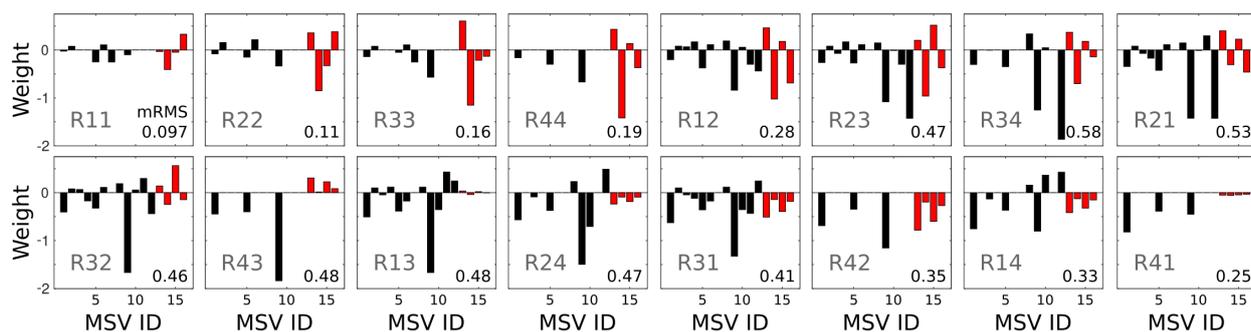


Figure B1. Model weights activated by an outlier on a single ray (e.g. R12, which connects shot 1 to receiver 2). The x-axis shows which model-space vector is activated, whereas the y-axis denotes how much it is activated. The red entries correspond to Null-Space vectors. These red values correspond to how much the corresponding data-space vector is activated. The number in the corner of each subplot shows the resulting model RMS misfit.

produce a single traveltimes, each outlier is reproduced by a linear combination of several DSVs (thus activating the corresponding MSVs). Fig. B1 shows that the outliers preferentially activate MSV 9 and MSV 12, similarly to what we saw in the Fig. 14.

The resulting models presented in Fig. 13 are summarized as model RMS misfits in Fig. B1. (As there is no true data, the results should be zeros everywhere.) These mRMS values again show, that individual outliers on the vertical rays (11, 22, 33, and 44) have the smallest impact on the final result. This is followed by the steepest diagonal rays (14 and 41).

Fig. B2 shows that the sensitivity of the individual MSVs to single outliers depends on which ray is affected. Interestingly, the influence of the outlier does not scale with the singular value, as per the common rule for errors (e.g., Menke 2018). MSV 7 is only affected with outliers on rays 1 and 3, whereas e.g. MSV 5 is always affected. Here also, MSVs 9 and 12 show the largest weights.

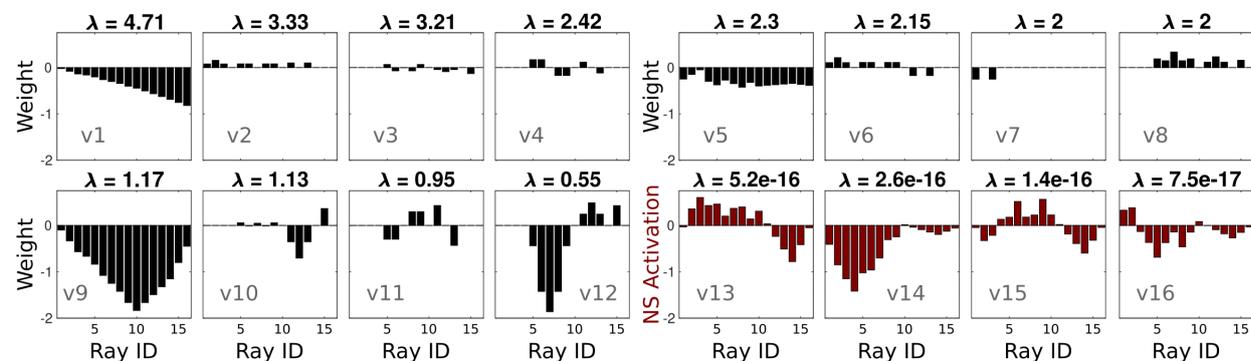


Figure B2. Model-Space vector (MSV) activation by an outlier on a single ray (Ray order according to Fig. B1 from left to right). The x-axis shows which ray has the outlier, whereas the y-axis denotes how much each MSV (e.g. v1) is activated. For the red entries (v13-v16), the length of the outlier in the direction of the data-space vector is depicted instead, as these vectors lie in the Null Space.

1091 We could not think of a clear reason why MSV 9 and MSV 12 should preferentially be activated
 1092 by such outliers. Nor, do we see a clear behavior here, which we could point out. The main point
 1093 of presenting these additional Figures, is that outliers do not behave the way they are commonly
 1094 explained (e.g. deviation in fitting a line, nor inverse proportional to the singular value).

1095 APPENDIX C: DECOUPLING STRATEGIES

1096 NS Transfer can be prevented by decoupling sub-components of MSVs in a geometry's NS. To
 1097 get a better understanding of what this means, we first need to dive into the decomposition of the
 1098 combined inversion problems. We will first present the examples for combined inversion methods
 1099 that produce multiple models, before explaining the alternate case for Joint inversion. For a method
 1100 that produces multiple models, when we apply SVD to both geometries simultaneously, without
 1101 applying any regularization (i.e. set $\alpha = 0$ in Eq. 7), this solves the two problems separately as we
 1102 can see in Eq. C.1 and C.2. Each column of $U_{no-regularization}$ and $V_{no-regularization}$ either solves for
 1103 the first, or for the second geometry (hence the two rows in each equation).

$$V_{no-regularization} = \begin{pmatrix} \overrightarrow{v_{11a}} & \vec{0} & \overrightarrow{v_{12a}} & \vec{0} & \cdots & \overrightarrow{v_{1ia}} & \vec{0} & |\overrightarrow{v_{1NS}} & \vec{0} & \cdots \\ \vec{0} & \overrightarrow{v_{21a}} & \vec{0} & \overrightarrow{v_{22a}} & \cdots & \vec{0} & \overrightarrow{v_{2ia}} & \vec{0} & \overrightarrow{v_{2i+1a}} & |\overrightarrow{v_{2NS}} \end{pmatrix} \quad (C.1)$$

$$U_{no-regularization} = \begin{pmatrix} \overrightarrow{u_{11a}} & \vec{0} & \overrightarrow{u_{12a}} & \vec{0} & \cdots & \overrightarrow{u_{1ia}} & \vec{0} & |\overrightarrow{u_{1NS}} & \vec{0} & \cdots \\ \vec{0} & \overrightarrow{u_{21a}} & \vec{0} & \overrightarrow{u_{22a}} & \cdots & \vec{0} & \overrightarrow{u_{2ia}} & \vec{0} & \overrightarrow{u_{2i+1a}} & |\overrightarrow{u_{2NS}} \end{pmatrix} \quad (C.2)$$

1104 Here, \overrightarrow{v} , corresponds to a sub-vector of V , and \overrightarrow{u} , corresponds to a sub-vector of U . The
 1105 first index denotes which geometry this sub-vector maps onto, and the second index denotes the
 1106 singular value index for the single geometry case. The last entry denotes if this sub-vector is native
 1107 to the geometry of the first index (a), or if this sub-vector fills in the constraints associated with a
 1108 native vector in the other geometry (b, see e.g. Eq. C.3). $\overrightarrow{u_{24a}}$ thus denotes the fourth DSV mapped
 1109 onto the 2nd geometry dominated by that geometry. The vertical bar denotes the cutoff for the NS,
 1110 and $\overrightarrow{v_{1NS}}$, corresponds to a MSV for geometry 1 in the NS. Note that the IS of two geometries do

1111 not need to have the same amount of MSVs, nor DSVs (like is the case for the geometries in Fig.
1112 2).

1113 The solutions for the sub-problems become coupled when we apply a regularization (Eqs. C.3
1114 and C.4):

$$V_{combined} = \begin{pmatrix} \overrightarrow{v_{11a}} & \overrightarrow{v_{11b}} & \overrightarrow{v_{12a}} & \overrightarrow{v_{12b}} & \cdots & \overrightarrow{v_{1ia}} & \overrightarrow{v_{1ib}} & |\overrightarrow{v_{1NS}} & \cdots & \cdots \\ \overrightarrow{v_{21b}} & \overrightarrow{v_{21a}} & \overrightarrow{v_{22b}} & \overrightarrow{v_{22a}} & \cdots & \overrightarrow{v_{2ib}} & \overrightarrow{v_{2ia}} & \overrightarrow{v_{2i+1b}} & \overrightarrow{v_{2i+1a}} & |\overrightarrow{v_{2NS}} \end{pmatrix} \quad (\text{C.3})$$

$$U_{combined} = \begin{pmatrix} \overrightarrow{u_{11a}} & \overrightarrow{u_{11b}} & \overrightarrow{u_{12a}} & \overrightarrow{u_{12b}} & \cdots & \overrightarrow{u_{1ia}} & \overrightarrow{u_{1ib}} & |\overrightarrow{u_{1NS}} & \cdots & \cdots \\ \overrightarrow{u_{21b}} & \overrightarrow{u_{21a}} & \overrightarrow{u_{22b}} & \overrightarrow{u_{22a}} & \cdots & \overrightarrow{u_{2ib}} & \overrightarrow{u_{2ia}} & \overrightarrow{u_{2i+1b}} & \overrightarrow{u_{2i+1a}} & |\overrightarrow{u_{2NS}} \\ \overrightarrow{l_1} & \overrightarrow{l_2} & \overrightarrow{l_3} & \overrightarrow{l_4} & \cdots & \overrightarrow{l_j} & \overrightarrow{l_{j+1}} & \cdots & \cdots & \cdots \end{pmatrix} \quad (\text{C.4})$$

1115 In Eq. C.3, the $\overrightarrow{0}$ below $\overrightarrow{v_{11a}}$ has been replaced with $\overrightarrow{v_{21b}}$. Thus, the corresponding singu-
1116 lar value is now shared by two sub-vectors. One that maps onto geometry 1 and one that maps
1117 onto geometry 2. This means that, whenever $\overrightarrow{v_{11a}}$ is included in the first model, $\overrightarrow{v_{21b}}$ is included
1118 in the second model. Eq. C.4 shows a similar adjustment, where the new sub-vector $\overrightarrow{u_{21b}}$ is the
1119 normalized data for the sub-vector $\overrightarrow{v_{21b}}$. $U_{combined}$ also has a third sub-vector, $\overrightarrow{l_i}$, which contains
1120 the regularization values, similar to the Laplacian values seen in Fig. 7. Note that the order of the
1121 vectors depends on the new singular values. Thus, a vector dominated by geometry 2 could come
1122 before one dominated by geometry 1. Each column will always have an "a" and "b" pair, however.

1123 In Eqs. C.3, $\overrightarrow{v_{1NS}}$, again corresponds to a MSV for geometry 1 in the NS for single inversion
1124 case. We define it in this way, as these vectors take part in the contamination. In the combined in-
1125 version, the corresponding singular value is no longer below the cutoff (due to the regularization),
1126 and thus no longer identifiable as being in the NS of our observables. This larger singular value
1127 allows the NS Transfer to occur.

1128 C1 Decoupling Trough Truncation With Cleanup

1129 Because NS Transfer occurs when a MSV in a geometry's NS is added to that geometry's result,
 1130 setting just these vectors back to $\vec{0}$ goes a long way towards negating NS transfer. Eq. C.5 presents
 1131 this strategy for methods that produce multiple models:

$$V_{decoupled} = \left(\begin{array}{cccc|cccc|ccc} \vec{v}_{11a} & \vec{v}_{11b} & \vec{v}_{12a} & \vec{v}_{12b} & \cdots & \vec{v}_{1ia} & \vec{v}_{1ib} & \vec{v}_{1NS} & \vec{0} & \vec{0} & \vec{0} \\ \vec{v}_{21b} & \vec{v}_{21a} & \vec{v}_{22b} & \vec{v}_{22a} & \cdots & \vec{v}_{2ib} & \vec{v}_{2ia} & \vec{v}_{2i+1a} & \vec{v}_{2i+1b} & \vec{v}_{2NS} & \vec{0} \end{array} \right) \quad (C.5)$$

1132 When vectors of both geometries are in their respective NS for the individual inversions, these
 1133 vectors can also be truncated completely. Because the singular values have been elevated, using
 1134 them as truncation criteria for this purpose will not ensure all NS components have been decou-
 1135 pled. As changing the problem also reorients the MSVs (as we have seen in Section 3.1.4), the
 1136 truncation strategy is unable to fully remove the coupling of all NS components. This is especially
 1137 evident for high α values (Fig. 18).

1138 C2 Decoupling By Purging Null-Space Entries

1139 In this strategy, we completely remove the components that lie in the NS for the single inversions
 1140 from $V_{combined}$. This is formalized in Eqs. C.6 and C.7:

$$v_{1,i,j}^{decoupled} = v_{1,i,j} - \sum_{k=1}^m \frac{v_k^{NS1} \cdot v_{1,i,j}}{v_k^{NS1} \cdot v_k^{NS1}} \cdot v_k^{NS1} \quad (C.6)$$

$$v_{2,i,j}^{decoupled} = v_{2,i,j} - \sum_{k=1}^n \frac{v_k^{NS2} \cdot v_{2,i,j}}{v_k^{NS2} \cdot v_k^{NS2}} \cdot v_k^{NS2} \quad (C.7)$$

1141 Here, v_k^{NS1} , are the MSVs in the NS of the single inversion for the first geometry, and, $v_{1,i,j}$ is
 1142 the sub-vector of $V_{combined}$ (e.g. \vec{V}_{11a}) corresponding to geometry 1. The sums use m and n ,
 1143 respectively, to signify that the geometries could have different numbers of MSVs in the NS.

1144 Because this strategy only changes the V matrix, the model weights of the combined problem
 1145 are not changed. Instead, the corresponding MSVs are included equal to the coupled problem,

1146 minus the NS entries. Thus the estimated models produced will not have the NS Transfer occur,
1147 and will otherwise be the same as the original combined problem.

1148 The purging strategy is not very practical, as it requires all NS entries. This is addressed in the
1149 third strategy, which produces equivalent results up to machine precision.

1150 **C3 Decoupling By Retaining Image-Space Entries**

One can get the same $v^{decoupled}$ by summing the components in the IS of the single inversions. This is formalized in Eqs. C.8 and C.9:

$$v_{1,i,j}^{decoupled} = \sum_{k=1}^m \frac{v_k^{IS1} \cdot v_{1,i,j}}{v_k^{NS1} \cdot v_k^{NS1}} \cdot v_k^{NS1} \quad (\text{C.8})$$

$$v_{2,i,j}^{decoupled} = \sum_{k=1}^n \frac{v_k^{IS2} \cdot v_{2,i,j}}{v_k^{NS2} \cdot v_k^{NS2}} \cdot v_k^{NS2} \quad (\text{C.9})$$

1151 This only requires the MSVs of the single inversions, making this strategy much more tractable.

1152 **C4 Decoupling NS Entries in Joint inversion**

1153 Joint inversion differs in two ways: The regularization is applied to the equations of one geometry,
1154 and only a single model is produced. Thus, no additional entry for the constraints is added to the
1155 DSVs (l_i in Eq. C.4), and the model-space vectors are not subdivided into an "a" and "b" pair.
1156 Therefore, the three strategies need to be implemented differently. For Joint inversion, all strate-
1157 gies need a decision on which NS components are to be decoupled: those of both geometries, or
1158 just of one. This choice will likely depend on the range of regularization values relevant to the
1159 problem. Both the purging strategy and the IS retention strategy then adjust the V matrix accord-
1160 ingly. (There will be only 1 sub-vector, corresponding to one model, so the chosen components of
1161 both geometries are applied to this one sub-vector.) Note, that adjusting the components of only
1162 one geometry is also possible for combined inversion methods that produce multiple models. In
1163 Joint inversion, this may make more sense to do, whereas more insights are gained for multi-model
1164 producing methods by adjusting the components of both geometries.

1165 Eqs C.10 - C.12 formalizes the NS purging strategy for Joint inversion, and Eqs. C.13 - C.15
 1166 formalize the IS retention for Joint inversion.

$$v_i^{decoupled,1} = v_i - \sum_{k=1}^m \frac{v_k^{NS1} \cdot v_i}{v_k^{NS1} \cdot v_k^{NS1}} \cdot v_k^{NS1} \quad (\text{C.10})$$

$$v_i^{decoupled,2} = v_i - \sum_{k=1}^n \frac{v_k^{NS2} \cdot v_i}{v_k^{NS2} \cdot v_k^{NS2}} \cdot v_k^{NS2} \quad (\text{C.11})$$

$$v_i^{decoupled,both} = v_i^{decoupled,1} - \sum_{k=1}^m \frac{v_k^{NS2} \cdot v_i^{decoupled,1}}{v_k^{NS2} \cdot v_k^{NS2}} \cdot v_k^{NS2} \quad (\text{C.12})$$

$$v_i^{decoupled,1} = \sum_{k=1}^m \frac{v_k^{IS1} \cdot v_i}{v_k^{NS1} \cdot v_k^{NS1}} \cdot v_k^{NS1} \quad (\text{C.13})$$

$$v_i^{decoupled,2} = \sum_{k=1}^n \frac{v_k^{IS2} \cdot v_i}{v_k^{NS2} \cdot v_k^{NS2}} \cdot v_k^{NS2} \quad (\text{C.14})$$

$$v_i^{decoupled,both} = \sum_{k=1}^n \frac{v_k^{IS2} \cdot v_i^{decoupled,1}}{v_k^{NS2} \cdot v_k^{NS2}} \cdot v_k^{NS2} \quad (\text{C.15})$$

1167 The truncation strategy makes less sense for Joint inversion. This method causes the following
 1168 two issues to arise when trying to implement this strategy: 1) When α values are low, the singular
 1169 values are dominated by one geometry. The singular values thus do not provide information on
 1170 where the components of the other geometry are located. Purging or retaining such components
 1171 will therefore not be straightforward. It will essentially require the same work as the NS purging
 1172 and IS retention strategies, with lesser performance (see the *Truncated* epoch-damping results in
 1173 Fig. 18). 2) The regularization can cause the order of the vectors to change relative to the single
 1174 inversions. Therefore, starting the truncation at the same vector index may not lead to remov-
 1175 ing the undesired components, and some desired components may be removed instead. Thus, the
 1176 truncation method likely produces poorer performance in Joint inversion, compared to the same
 1177 strategy applied to multi-model producing methods. We even expect poorer performance when
 1178 only adjusting the dominant geometry's components at low α values.

1179 **C5 What About Adjusting S and U?**

The NS purging and IS retention strategies change the MSVs in the V matrix, without making any adjustments to the singular values in the S matrix, nor the DSVs in the U matrix. Leaving the S and U matrices untouched equates to forcing a different answer to the same question. We have investigated what happens when recalculating S and U based on the new V matrix: $V_{decoupled}$. Here, we first normalize the MSVs in $V_{decoupled}$, which is not done in the strategies described above. Next, we calculate the synthetic data produced by $V_{decoupled}^{normalized}$, by multiplying it with the appropriate kernel (left-hand-side kernel in either Eq. 6 or 7). When α is non-zero, this will also produce the constraint values when using Eq. 7. This synthetic data is then normalized again, to produce the DSVs for the new U matrix: $U_{decoupled}$. The factors required to normalize the individual DSVs correspond to the inverse of the new singular values, which go into $S_{decoupled}$. Eqs. C.16 - C.18 formalize the creation and normalization of the U, S, and V matrices:

$$v_i^{normalized} = \frac{v_i}{v_i \cdot v_i} \quad (C.16)$$

$$\lambda_i = u_i \cdot u_i \quad (C.17)$$

$$u_i^{normalized} = \frac{u_i}{\lambda_i} \quad (C.18)$$

1180 The new set of matrices asks a different question of the data, and thus produces different
 1181 results, compared to the NS purging and IS retention strategies. Although the normalized versions
 1182 of these two strategies successfully prevent NS Transfer, the results sometimes differ strongly
 1183 from the previous strategies, resulting in a poorer performance. Thus we recommend the previous
 1184 strategies over setting up a self-consistent question (i.e. matching U, S, and V matrices).

1185 **C6 Unknown Consequences**

1186 These decoupling strategies need further investigation. It is, for example, not clear to us how NS
 1187 purging and IS retention needs to be implemented (and what the consequences would be) when
 1188 smoothing regularization is applied. An optimal integration of these strategies into a non-linear
 1189 combined inversion will also need to be found.