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# Ascent rates of 3D fractures driven by a finite batch of buoyant fluid

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(Received xx; revised xx; accepted xx)

Propagation of fluid-filled fractures by fluid buoyancy is important in a variety of settings, from magmatic dykes and veins to water-filled crevasses in glaciers. Industrial hydrofracturing utilises fluid-driven fractures to increase the permeability of rock formations; the effect of buoyancy on fracture pathways in this context is typically neglected. Analytical approximations for the buoyant ascent rate facilitate quantitative estimates of buoyant effects. Such analysis exists for two-dimensional fractures, but real fractures are 3D. Here we present novel analysis to predict the buoyant ascent speed of 3D fractures containing a fixed-volume batch of fluid. We provide two estimates of the ascent rate: an upper limit applicable at early time, and an asymptotic estimate (proportional to  $t^{-2/3}$ ) describing how the speed decays at late time. We infer and verify these predictions by comparison with numerical experiments across a range of scales and analogue experiments on liquid oil in solid gelatin. We find the ascent speed is a function of the fluid volume, density, viscosity and the elastic parameters of the host medium. Our approximate solutions can predict the ascent rate of fluid-driven fractures across a broad parameter space, including cases of water injection in shale and magmatic dykes. Our results demonstrate that both dykes and industrial hydro-fractures can ascend by buoyancy over a kilometre within a day.

**MSC Codes** 74R10: Brittle fracture 76D08: Lubrication theory

## 1. Introduction

### 1.1. Motivation

Fluid-filled fractures that propagate due to the buoyancy (positive or negative) of the fluid are found in a many geological settings. These fractures are discussed in the literature under various names including gravity-driven hydraulic fractures (e.g Salimzadeh *et al.* 2020), buoyant, liquid-filled cracks (e.g Dahm 2000; Taisne & Tait 2009) and fluid-filled fractures under the influence of weight contrasts (e.g. Pollard & Townsend 2018). Examples of this phenomenon include dykes and veins in the crust, as well as melt-water-filled crevasses in

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36 glaciers. Understanding the rate and extent of propagation of such fractures is critical for  
37 making useful, quantitative predictions. For example, the speed with which a batch of magma  
38 rises through the crust determines the duration of the warning period between geophysical  
39 detection of a dyking event and the volcanic eruption. Fracture ascent speed and distance  
40 also have implications for industrial operations. They determine how quickly buoyant fluid  
41 can travel away from an injection site, causing potential leakage into critically stressed faults  
42 or contaminating the groundwater in overlying geological strata.

43 Field and laboratory observations of fluid-driven fracture provide some indication of  
44 typical rates of ascent; for example, in basaltic systems this on the order of mm/s to  $\sim 0.5$  m/s  
45 (Mutch *et al.* 2019; Tolstoy *et al.* 2006). Observations of supraglacial lakes suddenly draining  
46 through growing crevasses indicate similar propagation rates (Das *et al.* 2008; Stevens *et al.*  
47 2015). Although there is good evidence that water- and gas-filled fractures can ascend through  
48 the crust (Schultz 2016; Cartwright *et al.* 2021), this process has not been observed directly.  
49 Estimates based on geochemical analysis give ascent rates of  $\sim 0.01$ – $0.1$  m/s, (1 km/day)  
50 (Okamoto & Tsuchiya 2009). Despite these estimates and their relevance for industrial  
51 fracking operations, there is a prevailing assumption that buoyant ascent of fluid-filled  
52 fractures is negligible or non-existent. This has a direct influence on assessments of the  
53 safety of such operations; for examples of this assumption influencing decision-making  
54 processes, see the UK and US scientific reviews of hydraulic fracturing safety (Mair *et al.*  
55 2012; EPA 2016).

## 56 1.2. *Previous work and current aims*

57 Analytical predictions of the speed and height of a fracture have previously been obtained  
58 from two-dimensional (2D) models. 2D models assume plane-strain elasticity, where there  
59 is zero strain out of the plane of interest ( $x$ – $z$ ) such that the fracture has an infinite extent  
60 with invariance along its third dimension ( $y$ ). 2D models cannot capture effects of the finite  
61 lateral extent in  $y$  of all real fractures. Nonetheless they provide useful physical insight and  
62 can be accurate along a central axis if the fracture is sufficiently broad in the  $y$  direction.

63 Two classes of 2D buoyant fractures have been characterised analytically. These are  
64 categorised according to the fluid source, either constant area-injection rate or a fixed finite  
65 area, where area refers to the volume per unit length in  $y$ . Early workers obtained solutions  
66 to the problem of a constant injection rate (e.g., Lister & Kerr 1991; Roper & Lister 2007;  
67 Rivalta *et al.* 2015). In these studies, injection is along a line-source of infinite length, at a  
68 rate that has units of volume per second per distance along the source,  $\text{m}^2 \text{s}^{-1}$ . At later times,  
69 the fracture head ascends buoyantly above a constant-aperture tail, which transmits a steady  
70 supply of fluid upward to the head.

71 In the other fluid-source class, Spence & Turcotte (1990) estimated the time-dependent  
72 ascent speed of a fixed, finite-area batch of fluid. Unlike the case of constant injection rate,  
73 this model predicts a diminishing aperture of the fracture tail, slowing fluid supply to the  
74 head, and hence causing the rate of propagation to diminish with time and ascent distance.

75 Analytical approximations for both of these classes have been shown to be consistent with  
76 2D numerical solutions (Lister 1990; Roper & Lister 2007; Dontsov & Peirce 2015). The  
77 numerics confirm that the elastic constants and material toughness influence the shape of the  
78 propagating head but not the ascent speed. The latter is determined by the width of the tail.

79 Three-dimensional (3D) fractures, with finite extent in the  $y$  direction, pose additional  
80 challenges to analysis. Workers have begun to investigate the case of constant volume injection  
81 rate in 3D (Möri & Lecampion 2021*b*), proving approximate analytical solutions for the height  
82 as a function of time, the width and the aperture of the fracture. In contrast, there has  
83 been no analysis of the ascent speed of a 3D fracture containing a fixed volume of fluid. And  
84 yet such analysis would be valuable in various contexts. In volcanology, where the volume

85 of intrusions is often well-constrained by geodetic data, a 3D solution would help to estimate  
 86 parameters that drive dyke ascent through the crust; it would enable rapid assessment of the  
 87 likelihood of eruption. In industrial operations, both the source rock properties and volumes  
 88 of fluid injected are well constrained, but currently there is no simple way of predicting how  
 89 fast an injected batch of fluid will ascend by buoyancy.

90 Here we develop an analytical solution that predicts the size and ascent speed of a 3D  
 91 fracture driven by a finite volume of buoyant fluid. In developing this solution, our first aim  
 92 is to provide an upper bound on the ascent speed of fractures that are propagating due to the  
 93 buoyancy of a finite batch of fluid. Our second aim is to understand how this ascent speed  
 94 decays with time, as in the 2D solution of Spence & Turcotte (1990).

95 Our strategy is to use a state-of-the-art numerical simulator (Zia & Lecampion 2020) to  
 96 produce 3D solutions to the full, non-linear equations. We treat these as a benchmark for our  
 97 novel analytical results, to show their applicability. The manuscript is organised as follows.  
 98 In the next section we review existing analytical approximations and introduce the simulator  
 99 of Zia & Lecampion (2020). We use this to simulate the ascent of a fixed finite-volume fluid  
 100 batch, reviewing this simulation in detail to describe the problem. In section 3, we use these  
 101 insights and proceed to derive equations describing the ascent speed of a 3D, finite batch of  
 102 fluid. In section 4, we use 3D simulators to test the validity of these results across a range of  
 103 scales. We then test the ability of our approximations to predict the ascent speed of oil-filled  
 104 cracks in gelatin solids. Lastly, in Sec. 5, we discuss the implications of our results in relation  
 105 to selected case studies, detailing limitations and potential avenues of future research.

## 106 2. Background

107 In this section we first review approximate analytical solutions for two- and three-dimensional  
 108 fractures, then consider numerical simulators that solve the full, non-linear problem.

### 109 2.1. Analytical solutions

#### 110 2.1.1. Critical lengths and volumes

111 A fluid-filled fracture will either ascend or descend, depending on the density of the fluid  
 112 relative to that of the surrounding solid medium. This has been well documented in analog  
 113 experiments by comparing, for example, the injection of air and mercury into gelatin solids  
 114 as described in Heimpel & Olson (1994). For simplicity here, we consider only positively  
 115 buoyant (less dense) fluids, and hence we model only fracture ascent, noting that there is no  
 116 loss of generality in the solutions derived.

117 Fluid-filled fractures are driven to ascend by a weight contrast  $\Delta\gamma$  between the fluid and  
 118 surrounding rock. More specifically, the buoyancy force is defined as the difference between  
 119 the vertical gradient of the horizontal stress in the rock column and the volume-specific  
 120 weight of the magma contained within a vertical fracture. Assuming the horizontal and  
 121 vertical stresses are equal, i.e., a lithostatic state of stress, then  $\Delta\gamma = (\rho_r - \rho_f)g$  (Secor  
 122 & Pollard 1975), where  $\rho_r$  and  $\rho_f$  are rock and fluid density, respectively, and  $g$  is the  
 123 acceleration due to gravity.

124 Weertman (1971) and Secor & Pollard (1975) showed that fluid-filled fractures can ascend  
 125 through their solid host by hydraulic fracturing, provided a critical areal extent  $A_c$  is exceeded,

$$126 \quad A_c = \frac{(1 - \nu) K_c^2}{2\mu \Delta\gamma}. \quad (2.1)$$

127 Here the stress intensity factor  $K_I$  appears as  $K_c$  to represent the fracture toughness of the  
 128 host rock for mode-I fracture (Tada *et al.* 2000). This is derived by requiring that the fracture

129 walls close shut at the lower tip ( $K_I(z = -c) = 0$ ) and that the upper tip is critically stressed  
 130 ( $K_I(z = +c) = K_c$ ), where  $c$  is the fracture's half-length and  $z$  the vertical distance from  
 131 its centre (Pollard & Townsend 2018). Working independently, Dahm (2000); Salimzadeh  
 132 *et al.* (2020); Davis *et al.* (2020); Smittarello *et al.* (2021); Möri & Lecampion (2021*b*) each  
 133 extended the 2D analytical model of Secor & Pollard (1975) to quantify the critical volume  
 134 of fluid for ascent of a 3D fracture. The solutions derived in these studies have the same  
 135 scaling with parameters but differ by around 10% because of different numerical constants.  
 136 The solution for critical volume by Davis *et al.* (2020) is

$$137 \quad V_c = \frac{(1 - \nu)}{16\mu} \left( \frac{9\pi^4 K_c^8}{\Delta\gamma^5} \right)^{1/3}, \quad (2.2)$$

138 where  $\nu$  and  $\mu$  are Poisson's ratio and the shear modulus of the host rock, respectively. When  
 139 the volume of fluid in a fracture exceeds  $V_c$ , the ascent of the fracture and contained fluid  
 140 is self-sustaining. By this we mean that the ascent is due to buoyancy alone, and requires  
 141 no additional forces such as a driving pressure (e.g., from a pressurised magma chamber or  
 142 well bore). As the fracture ascends, the tail of the crack lengthens and thins, but the fracture  
 143 head retains a volume sufficient to critically stress the medium ahead of it. Figure 1 shows  
 144 numerical solutions, discussed below, that illustrate this.

### 145 2.1.2. Ascent rate

146 Results cited above quantify the critical conditions for buoyant ascent, but not the speed of  
 147 that ascent. The latter requires a consideration of the fluid flow within the fracture. In the tail  
 148 of the fracture this can be usefully approximated as Poiseuille flow between parallel plates.  
 149 At small Reynolds number, lubrication theory gives the mean flow speed  $v$  and areal flux  $Q$   
 150 as

$$151 \quad v = \frac{w^2}{12\eta} \Delta\gamma, \quad Q = \frac{w^3}{12\eta} \Delta\gamma, \quad (2.3)$$

152 where  $w$  the separation of the plates,  $\eta$  is the dynamic viscosity and we have assumed that  
 153 the gravity vector is parallel to the walls. Here it is assumed that the fluid flow between the  
 154 fracture's walls is laminar.

155 The ascent rate of a 2D fracture with constant areal flux  $Q$  is given by eliminating  $w$  from  
 156 equations (2.3) to give (Lister & Kerr 1991)

$$157 \quad v_Q = \left( \frac{Q^2 \Delta\gamma}{12\eta} \right)^{1/3}. \quad (2.4)$$

158 This result means that for  $Q$  constant, the speed of the upper tip is controlled by viscous flow  
 159 through a tail of constant width, which transports buoyant fluid upward to the propagating  
 160 head.

161 The ascent rate of a 2D fracture with constant area  $A$  is given by combining (2.3) with a  
 162 conservation of mass equation that relates the rate of change of aperture to the divergence of  
 163 the vertical flux (Spence & Turcotte 1990). On this basis, Spence & Turcotte (1990) obtained  
 164 a solution in terms of the velocity of flow through a half-ellipse with a fixed area. As this  
 165 lengthens vertically its aperture decreases, which hinders fluid flow and slows the ascent of  
 166 the fracture tip. The rate at which the half-ellipse lengthens therefore progressively decreases  
 167 with time. Differentiating with respect to time equation (22) of Spence & Turcotte (1990),  
 168 which defines the  $z$  distance between the upper tip and the initial centre-point of the fracture

169 (the injection point), the fracture’s tip velocity is

$$170 \quad v_A(t) = \left( \frac{A^2 \Delta \gamma}{48 \eta} \right)^{1/3} t^{-2/3}. \quad (2.5)$$

171 The ascent rate of a 3D fracture with a constant rate of volume injection was considered  
 172 by Lister (1990). His equation (2.14) provides an approximate 3D scaling for fracture height,  
 173 breadth and aperture at a given injection rate (see also Germanovich *et al.* 2014). These  
 174 scalings have been confirmed by Möri *et al.* (2020) and Möri & Lecampion (2021*b*) using  
 175 3D numerical simulations computed with open-source codes (Zia & Lecampion 2020;  
 176 Salimzadeh *et al.* 2020). Möri *et al.* (2020) and Möri & Lecampion (2021*b*) show that  
 177 for a 3D, constant injection-rate, ascending fracture, the shape of the tip-line depends  
 178 on a dimensionless ratio that includes viscosity and fracture toughness. Close to the  
 179 injection point, viscous resistance limits tip-line propagation, resulting in a V-shaped tip  
 180 line; sufficiently far from the injection point, fracture toughness limits tip-line propagation  
 181 and the tip-line at either side of the fracture is vertical. These limiting regimes correspond  
 182 with the asymptotic solutions for aperture near the tip-line (Detournay 2004).

183 The ascent rate of a 3D fracture containing a constant volume of fluid is not known.

## 184 *2.2. Numerical Methods*

185 We run 3D simulations of an ascending fluid batch with a constant-volume using a  
 186 hydro-fracture simulator. This 3D simulator generates numerical solutions to the nonlinear  
 187 governing equations. These solutions guide the development of our approximate analysis  
 188 for constant-volume fractures. Comparison of our analytical and numerical solutions then  
 189 enables us to determine the validity of simplifying assumptions and assess the accuracy of  
 190 our analytical predictions.

191 We simulate the propagation of buoyancy-driven fractures using the open source code  
 192 PyFrac (Version 1.0, <https://pyfrac.epfl.ch>). PyFrac’s methods and implementation  
 193 are extensively documented in Zia & Lecampion (2020). The calculations we present use the  
 194 code as described in that original work, and hence we only summarise the algorithm here to  
 195 give a sense of its aptitude for solving hydro-fracture problems.

196 The hydro-fracture problem comprises three coupled physical aspects: fluid flow and  
 197 pressure inside the fracture, deformation of the surrounding medium due to the fracture,  
 198 and propagation of the fracture’s tip-line into the unfractured medium. These aspects are  
 199 independently modelled as low Reynolds-number fluid flow, quasi-static linear elasticity, and  
 200 linear-elastic fracture mechanics, respectively. Although each is linear independently, their  
 201 coupling results in a non-linear equation system. The time-dependent solution to this system  
 202 is the location of the tip-line, the fracture aperture, and the fluid pressure and velocity. The  
 203 PyFrac solver limits the complexity of this problem by restricting solutions to be planar  
 204 fractures. This simplification enables pre-computation of the Green’s functions required  
 205 for the Displacement Discontinuity Method (DDM) using a fixed-grid mesh (Peirce 2016).  
 206 The DDM method links fluid pressure inside the fracture to the opening of its walls and  
 207 to deformation of the surrounding medium. On the same mesh, the lubrication equation is  
 208 discretised using the finite volume method at element centres. The gradients in fluid pressure  
 209 determine the rate of flow between cells. Crucially, to avoid the requirement of ultra-high mesh  
 210 resolution near the fracture tip, the PyFrac uses near-tip asymptotic solutions for propagating,  
 211 fluid-filled fractures. These asymptotic solutions resolve the sub-grid-scale dynamics, and  
 212 hence are used to assess the propagation criterion and rate of tip-line motion. By comparison  
 213 of numerical solutions to existing analytical solutions (Peirce 2015, 2016; Zia *et al.* 2018;

214 Zia & Lecampion 2020; Moukhtari *et al.* 2020), it has been shown that PyFrac can accurately  
215 solve the coupled hydro-fracture growth problem.

216 For simulations in the present manuscript the extent of the meshed domain in the horizontal  
217 direction is  $2a$ , the characteristic radius we define below. The vertical extent of the domain is  
218 sufficient to capture propagation of the fracture in the post-injection, buoyancy-driven regime  
219 without re-meshing (14a). Along the horizontal direction we use a minimum grid size of 50  
220 elements, the vertical grid is created such that the elements are square. We remind readers  
221 the meshed domain is only used to discretise the fracture itself. Displacements and stresses  
222 due to the fracture are described by Green's functions for an infinite space such that these  
223 hold anywhere in the body. Injection of the fluid occurs from a point source.

### 224 2.3. *Qualitative insights from numerical simulations of fracture ascent*

225 Using the numerical method described above, we simulate the injection of a finite batch of  
226 buoyant fluid, closely following a case that was analysed by Salimzadeh *et al.* (2020). The  
227 parameter values for this simulation are provided in Table 1.

228 In Fig. 1 we plot snapshots of the fracture's shape as it evolves through time. Time  
229 increases from left to right in the plot. Looking onto the fracture's face, soon after the  
230 injection has ceased (10 minutes), the fracture's tip-line is circular. Note the asymmetric  
231 aperture of the cross-section at this time, which indicates that the fluid is starting to flow  
232 towards in the upper tip of the fracture. After 6 hours the top of the fracture has grown  
233 upwards whilst the tip-line in lower region has not changed shape; when looking at the  
234 face at this stage, the tip-line on either side of the fracture tapers progressively towards the  
235 top. The form of the upper part of the tip-line is a semi-circle and remains so throughout  
236 the rest of the simulation. As time increases the aperture in the lower parts of the fracture  
237 progressively thins. 6 hours into the simulation, the fracture has reached a characteristic  
238 shape; in cross-section, the lower half of the fracture (tail) is thin and V-shaped while the  
239 upper half (head) has formed the teardrop profile of a Weertman crack (Roper & Lister 2007).  
240 Looking onto to the face at 15 hours, the fracture's upper tip has continued to rise, resulting  
241 in a tip-line on either side of the fracture that is approximately vertical, where the tapering  
242 upwards is less evident. The cross-sectional aperture retains the typical structure predicted by  
243 2D analytical solutions, with a clearly defined teardrop-shaped head and tapered tail (Roper  
244 & Lister 2007). This structure dominates at late times ( $t > 15$  hours), noting the thinning of  
245 the tail as the length increases, whilst the head's shape remains roughly constant.

246 In Fig. 2a we plot the speed of the upper-most point on the fracture's tip-line as a function  
247 of time (log scale). We refer to this speed as the ascent rate because it is identical to the  
248 speed at which the head of the fracture rises through the medium. A dashed line showing  
249  $t^{-2/3}$  from equation (2.5) is overlain. At late-times  $> 6$  hrs, the velocity of the 3D numerical  
250 solution approaches the  $t^{-2/3}$  asymptote described by the 2D analytical solution.

251 In Figure 2b we plot the ascent rate as a function of the height of the fracture. Note the  
252 y-axis has a logarithmic scale and the times shown in Fig. 1 are plotted as asterisks. In the  
253 later stages of the simulation, the ascent rate has undulations that repeat every few iterations  
254 of the simulation; we interpret these as being due to numerical instability. The undulations  
255 have a roughly uniform magnitude that depends on the mesh size. Neglecting these, the plot  
256 shows two clear changes in front deceleration. The first occurs once the injection stops (Möri  
257 & Lecampion 2021a), and the second once the upper tip has propagated further than  $2a$ ,  
258 a characteristic radius we examine later. We interpret the second transition in the ascent  
259 speed to a shift in the process driving tip-line growth, from a radial injection-rate-driven  
260 fracture growth to buoyancy-driven fracturing, occurring once the upper tip of the fracture  
261 has propagated a given distance. If this interpretation is correct then at later times the decay  
262 in ascent speed should be related to the fluid draining upward from the increasingly long

Quantity	Symbol	Unit	Value
Young's modulus	$E$	Pa	$20 \times 10^9$
Poisson's ratio	$\nu$		0.25
Fracture toughness	$K_c$	$\text{Pa m}^{1/2}$	$2.0 \times 10^6$
Fluid/rock density difference	$\Delta\gamma$	$\text{kg m}^{-3}$	1000
Fluid viscosity	$\eta$	Pa s	0.05
Injected Volume	$V_I$	$\text{m}^3$	1.95
Critical Volume	$V_c$	$\text{m}^3$	0.79
Volume ratio	$V_I/V_c$		2.47
Injection rate	$Q(t)$	$\text{m}^3 \text{s}^{-1}$	0.015

Table 1: Parameter values used for the simulation shown in Figs. 1 & 2. These correspond to simulation-case II from the main text of Salimzadeh *et al.* (2020).

263 and thin tail (Eq.2.5) (Spence & Turcotte 1990). Such a thinning of the tail aperture can be  
 264 seen in the snapshots of the ascending fracture in Fig. 1. In the following section we provide  
 265 analytical expressions approximating: (i) the fracture's maximum lateral extent  $a$  (Fig. 1),  
 266 (ii) the maximal upwards propagation speed once injection has terminated (dashed line in  
 267 Fig. 2b)), and (iii) the deceleration of the ascending fracture with time (Fig. 2a).

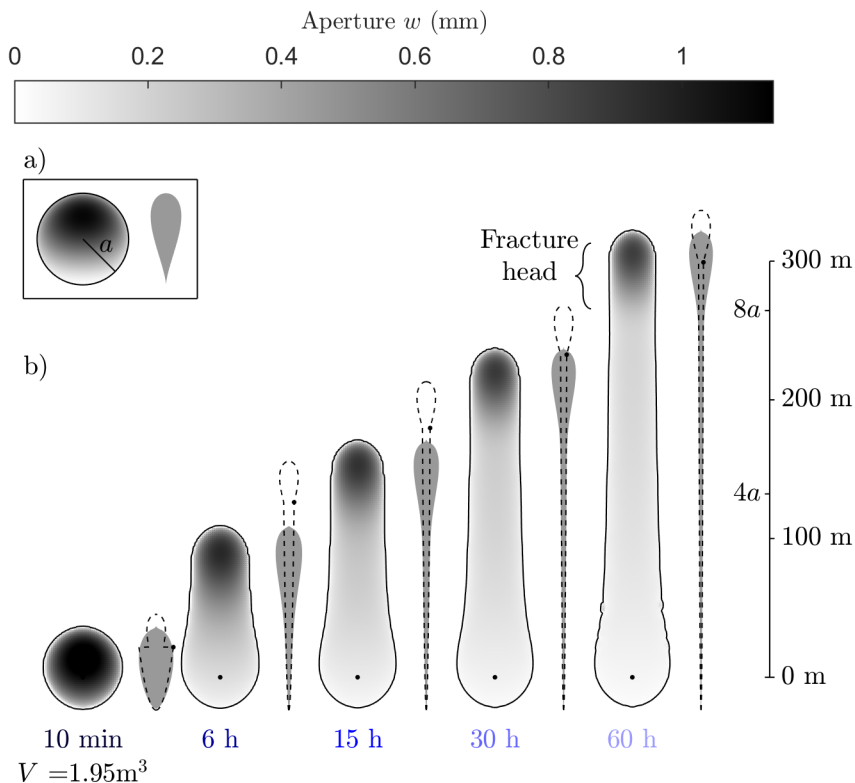


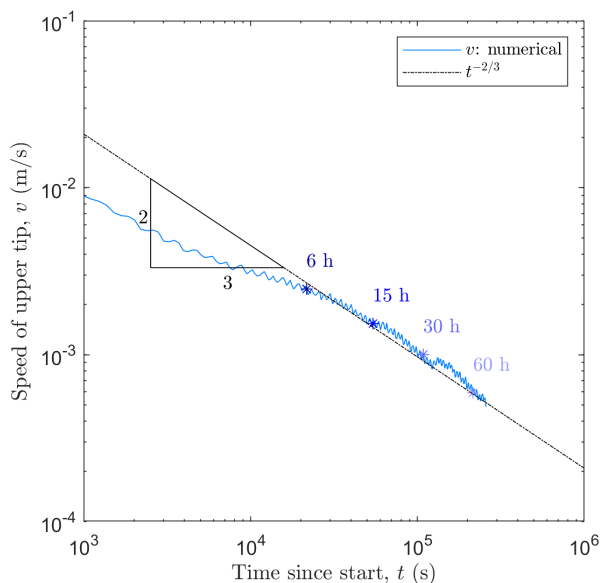
Figure 1: Numerical simulation of an ascending finite fluid batch. In this simulation  $1.95\text{ m}^3$  of fluid was injected over 130 seconds. Two views of the fracture are shown at each time. On the left is a view looking onto the fracture's face, shaded by aperture ( $w$ ) with the tip-line in black. On the right is a grey cross-section showing the profile of aperture along the centre of the crack. The latter has a horizontal exaggeration of  $2 \times 10^4$ .

a) Analytical solution that defines the radius  $a$ . The plotted solution contains fluid volume  $V$  and has a radius such that  $K_I(z = -a) = 0$ .

b) Numerical simulation of fracture ascent using PyFrac (Table. 1). Simulation times are shown below the respective fractures. The time dependant analytical approximation of the fracture's cross-section is shown as dashed lines (Roper & Lister 2007). The tail height  $h$  (3.5) is marked with a dot.



a)



b)

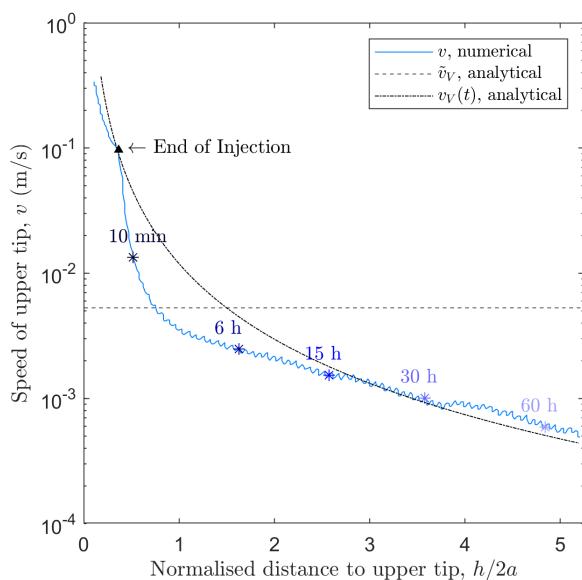


Figure 2: Upper tip ascent speed from PyFrac simulation (solid line) and analytical approximations (broken lines). Times corresponding to steps in Fig. 1 are shown as asterisks. Parameter values are given in table 1. *a*) Ascent rate versus time, log-log plot showing that at late-times the simulations approach the  $t^{-2/3}$  asymptote (dash-dot). *b*) Ascent rate versus height, showing our equation for the maximum ascent speed (3.4) as a horizontal dashed line, and the analytical approximation of the decelerating front speed produced using equation (3.5) for  $x$  values and (3.6) for  $y$  values (dash-dot line).

### 268 3. Analysis of three-dimensional fracture ascent rates

#### 269 3.1. Early-time ascent rate

270 As a first pass, we aim to approximate the ascent rate of a buoyancy-driven fracture at an  
 271 early time. Here we refer to the moment after injection when ascent becomes dominantly  
 272 driven by buoyancy. This moment does not occur precisely at the end of injection; some  
 273 ascent beyond that time is driven by the release of elastic energy stored during the injection  
 274 phase, which drives a radial motion of the tip line (Möri & Lecampion 2021a). This energy  
 275 is expended rapidly, after which buoyancy dominates in driving ascent. Our first calculation  
 276 is an estimate of the speed of ascent at this point.

277 We assume that a volume  $V$  of fluid has been injected and resulted in a penny-shaped crack.  
 278 The crack has extended radially to a size such that the deepest segment of the tip line, the  
 279 bottom tip, has ceased its radial advance. The walls of the crack are subject to a downward  
 280 stress gradient of magnitude  $\Delta\gamma$  that drives fluid upward. As this gradient becomes dominant  
 281 post-injection, the fracture above the bottom tip begins to drain fluid and pinch closed.

282 Injection of a volume  $V$  leads to a penny-shaped crack with mean internal pressure given  
 283 by

$$284 \quad p_0 = \frac{3\mu}{8(1-\nu)} \frac{V}{a^3}, \quad (3.1)$$

285 where  $a$  is the radius of the tip-line around the injection point (Tada *et al.* 2000). The mode-I  
 286 stress intensity at the circular tip line arises from a combination of  $p_0$  and the linear stress  
 287 gradient (Tada *et al.* 2000),

$$288 \quad K_I(\theta) = \frac{2}{\pi} \sqrt{\pi a} \left( p_0 + \frac{2}{3} \Delta\gamma a \cos(\theta) \right), \quad (3.2)$$

289 where  $\theta$  is the angle in the  $y$ - $z$  plane away from vertical. If  $z$  is the upward distance from the  
 290 injection point, eliminating the mean pressure from equations (3.1) and (3.2) and requiring  
 291 that at the basal tip  $K_I(z = -a) = 0$  gives the radius of the crack as

$$292 \quad a = \left( \frac{9\mu}{16(1-\nu)} \frac{V}{\Delta\gamma} \right)^{1/4}. \quad (3.3)$$

293 The mean aperture of this crack is simply given by  $w = V/\pi a^2$ ; using this in (2.3) to compute  
 294 the characteristic fluid ascent speed  $v$ , then combining the result with (3.3) to eliminate  $a$  we  
 295 obtain

$$296 \quad \tilde{v}_V = \frac{4(1-\nu)}{27\pi^2\mu} \frac{\Delta\gamma^2 V}{\eta} \quad \text{for } V > V_c. \quad (3.4)$$

297 Here we have assumed that the fluid ascent speed  $v$  is equal to the speed of the upper tip  
 298 line  $\tilde{v}_V$  (the tilde indicates that this is an early-time solution and the subscript  $V$  indicates  
 299 that this is a fixed-volume, 3D fracture). This speed is plotted as a horizontal dashed line in  
 300 Figure 2b.

301 Equation (3.4) is the speed of the upper tip when the crack is still penny-shaped, just after  
 302 the phase of radial growth driven by the injection. As the upper tip propagates driven by  
 303 buoyancy of the fluid, the crack elongates vertically. The stress gradient drives fluid within  
 304 the crack upward such that part of the total fluid volume  $V$  resides in the head the crack, while  
 305 the remainder forms a slowly thinning layer in the long tail (Fig. 1). Thus, equation (3.4),  
 306 which describes the instant when the entire fluid volume is in the head, gives an approximate  
 307 upper bound on the buoyancy-driven ascent speed.

308

### 3.2. Finite propagation with deceleration

309 We now consider the vertical advance of the fracture over finite time. We define  $h$   
 310 as the vertical distance between the fracture's upper tip and the initial centre of the  
 311 crack (the injection point). Recalling that Fig.2a shows that the ascent rate of the 3D  
 312 simulation asymptotically scales with time as predicted by equation (2.5), this motivates  
 313 us to approximate the 3D ascent speed using this 2D solution. To retrieve  $h(t)$  and  $v(t)$  for  
 314 our 3D fracture, we modify the similarity solution obtained by Spence & Turcotte (1990,  
 315 eq. 22) for a 2D fracture with constant area. Replacing that area  $A$  with the initial cross-  
 316 sectional area of the 3D crack  $2aw$  gives

$$317 \quad h = \left( \frac{9 (2aw)^2 \Delta\gamma t}{16\eta} \right)^{1/3}. \quad (3.5)$$

318 Substituting equation (3.3) and  $w = V/\pi a^2$  into (3.5) and defining the speed  $v_V(t) \equiv dh/dt$ ,

$$319 \quad v_V(t) = \left( \frac{(1-\nu) \Delta\gamma^3 V^3}{9^2 \pi^4 \mu} \frac{1}{\eta^2} \right)^{1/6} t^{-2/3}. \quad (3.6)$$

320 This approximate solution might be expected to hold at later times, when buoyancy has  
 321 driven the fracture to propagate away from the injection point, establishing distinct tail and  
 322 head regions of the fracture. Assuming that this structure is in place once the upper tip has  
 323 advanced to a height  $h = 2a$ , we use (3.5) to define late time as  $t > t_r$  where

$$324 \quad t_r = \left( \frac{3^2 \pi^8 \mu^5}{(1-\nu)^5} \frac{\eta^4}{\Delta\gamma^9 V^3} \right)^{1/4}. \quad (3.7)$$

325 Figure 2b shows the predicted initial and late-time speeds from equations (3.4) and (3.6),  
 326 respectively. The figure also shows a numerical solution of a 3D crack for the same parameter  
 327 values. The ascent rate predicted by (3.6) at  $h = 2a$  (corresponding to  $t = t_r$ ) is greater than  
 328 the numerical result at the same position; at this point, the early-time approximation (3.4) fits  
 329 better. At  $h/2a = 3/2$ , the two approximations are equal  $\tilde{v}_V = v_V(t)$ . As the fracture ascends  
 330 beyond this height, the full, numerical solution and the late-time solution converge. This  
 331 convergence was also obtained in 2D numerical solutions by Roper & Lister (2007, Fig. 9).

332

### 3.3. Comparison between the numerical results and analytical predictions

333 Using the initial cross-sectional area  $2aw$  noted above, we plot as dashed lines in Fig. 1  
 334 the cross-section predicted by the 2D similarity solutions of Roper & Lister (2007, Sec. 6)  
 335 (Also see our Appendix A). The height of the tail in these cross-sections is given by equation  
 336 (3.5). Focusing attention on the solutions at elapsed time of 30 hours, the analytical solution  
 337 provides a good fit to the width of the tail from the numerical solution. Furthermore, the  
 338 simulated fracture's head length and shape are approximately that of the Weertman solution.  
 339 The notable discrepancy is that analytical solution predicts a greater propagation distance than  
 340 produced by the simulation. This discrepancy is also evident in the comparison of fracture  
 341 ascent speeds of Fig. 2b where, initially, the analytical prediction is for faster propagation,  
 342 but as the simulation progresses and the fracture lengthens, the numerical and analytical  
 343 speeds converge. Next we discuss the assumptions made in the derivation of (3.6) to explain  
 344 why the numerical and analytical ascent rates and heights differ.

345 We begin with a reminder of three insights from 2D results by Roper & Lister (2007) that  
 346 also apply here. Firstly, as long as the cross-sectional area of the initial 2D crack is above  
 347 the critical head area given by equation (2.1), the fracture is predicted to ascend indefinitely,  
 348 with a monotonically decreasing propagation rate. Secondly, in the 2D, constant-area theory,

349 elastic forces dominate in the head region, resulting in a head shape that remains constant  
 350 over time and is described by the static Weertman solution (Weertman 1971; Rubin 1995).  
 351 Thirdly, Roper & Lister (2007) show that the area of the head should be removed from  
 352 the total area when computing the dynamics of the tail (e.g., (3.5)) and, in particular, in  
 353 estimating the velocity and height of the fracture.

354 Keeping these insights in mind, we next evaluate assumptions of the 2D, constant-area  
 355 approximate solution to observations from our 3D simulation from Fig. 1. In the 2D solution,  
 356 the areal extent of the head is constant during ascent, whereas in the 3D simulation, the  
 357 effective head volume decreases with time. This is qualitatively shown by the lateral tip-  
 358 lines in Fig. 1 that slightly converge upwards over time. Because of this feature, we have  
 359 neglected to approximate the head volume, placing the entire cross-sectional area of the  
 360 initial penny-shaped crack into the tail solution. Our analytical result is therefore based on  
 361 an overestimate of the tail volume, and hence it should predict a faster ascent rate than  
 362 the numerical simulation. This explains why initially, the predicted velocities are faster.  
 363 Moreover, in the 2D solution (3.5), the cross-sectional area of the fluid (tail plus head) is  
 364 constant. This contrasts with observations in the simulation shown in Fig. 1 where along the  
 365 centreline, the fracture's cross-sectional area increases by 36% from the earliest (10 minutes)  
 366 to the latest (60 hours) snapshots. This out-of-plane fluid flow is, by definition, not captured  
 367 by the 2D solution. Thickening of the cross-section in the 3D simulation suggests that the  
 368 ascent rate for a 3D fracture will be faster than its 2D analogue at later times.

### 369 3.3.1. *Turbulent flow*

370 Up to this point, we have assumed that the crack aperture is small and that the viscosity  
 371 is large enough that a lubrication approximation can be applied to the flow. To test this  
 372 assumption of laminar flow, the Reynolds number can be computed using

$$373 \quad \text{Re} = \frac{wv\rho_f}{\eta}, \quad (3.8)$$

374 where, as a first approximation, crack aperture  $w$  can be obtained according to  $w = V/\pi a^2$   
 375 and  $v$  can be taken as the initial buoyancy-driven ascent rate from equation (3.4). For flow  
 376 between parallel plates, Reynolds numbers larger than  $\sim 1400$  are associated with turbulence.  
 377 Here we avoid empirical formulations describing flow speeds in turbulent regimes for the  
 378 sake of simplicity, focusing instead on the upper limit of the ascent speed that is described  
 379 by laminar flow.

380 For analog experiments using air-filled cracks in gelatine, (e.g. Heimpel & Olson (1994),  
 381 Smittarello *et al.* (2021)), equation (3.4) vastly overestimates measurements of the crack  
 382 ascent rate. For example, assessing the Reynolds number for an air-filled crack in a gelatin  
 383 solid with a stiffness  $E$  of 1500 Pa, toughness  $K_c$  of 50 Pa m<sup>1/2</sup> and a volume calculated  
 384 from equation (2.2), we find that is greater than 2900, and hence turbulent flow is expected  
 385 and may account for the slow ascent.

Quantity	Symbol	Unit	Oil in gelatin	Water in shale	Basaltic dyke
Shear modulus	$\mu$	Pa	276	$8 \times 10^9$	$25 \times 10^9$
Poisson's ratio	$\nu$		0.25	0.25	0.25
Fracture toughness	$K_c$	Pa m <sup>1/2</sup>	19	$2.0 \times 10^6$	$6.0 \times 10^6$
Fluid/solid density difference	$\Delta\gamma$	kg m <sup>-3</sup>	160	2,000	50
Fluid viscosity	$\eta$	Pa s	$4.8 \times 10^{-2}$	$5 \times 10^{-3}$	20
Critical Volume (2.2)	$V_c$	m <sup>3</sup>	$1.31 \times 10^{-5}$	0.7921	700

Table 2: Properties used to numerically simulate fracture ascent for different physical processes. The injection rate for such cracks is set to  $(V/\bar{v}_V)/a$ , such that the injection is complete before the fracture ascends past  $h = 2a$ .

## 386 4. Validation of analytical approximations

### 387 4.1. Testing across length scales by comparison to numerical solutions

388 We now test our analytical approximations against numerical solutions across a wider range  
389 of physical parameters. Here, each parameter set represents a different physical context:  
390 analog experiments with oil-filled cracks in gelatin, an industrial setting where water is  
391 injected into a shale sequence and, lastly, the ascent of a basaltic dyke through the crust  
392 (Table 2). We compare our analytical prediction from equation (3.6) against rates obtained  
393 from PyFrac calculations using the three parametric combinations listed in Table 2, each of  
394 which is simulated for three different injection volumes to give nine total models. Results  
395 are shown in Figure 3, in comparison with analytically predicted rates. The comparison  
396 shows that  $v_V(t)$  from (3.6) provides a reasonable match to the numerical results across  
397 all parameter sets considered. The undulations in propagation speed from PyFrac are again  
398 clear in this plot, and again we attribute these to numerical instability. Ignoring them, the  
399 analytical results predict the PyFrac velocities within a factor of two. The ratio of numerical  
400 to analytical ascent speed is approximately constant for a given simulation. In Figure 4,  
401 we compare our analysis (3.5) to numerical results in terms of the time-dependent height  
402 of the upper tip above the injection point. As in the previous figures, we find the fracture  
403 must ascend by a length of more than  $2a$  (box shown is  $3a$ ) before buoyancy becomes the  
404 dominant driver of propagation. After that time, the results show the predicted and simulated  
405 heights are comparable. These comparisons show that our analytical approximations capture  
406 the leading-order fracture ascent speed across a broad parameter space. We further note that  
407 our analytical results for the extent and ascent speed provide a reliable means to forecast the  
408 required PyFrac domain size and simulation duration for a given set of parameters, such that  
409 the numerical model converges and completes the intended simulation.

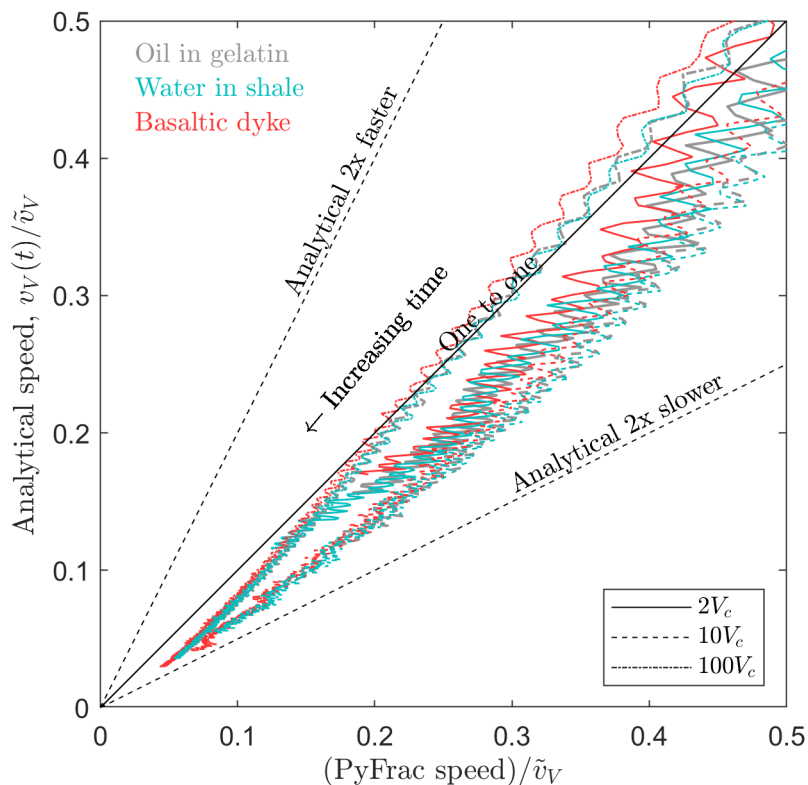


Figure 3: Numerical versus analytical (3.6) speed estimates at times  $t$  since injection. Plotted are nine simulations performed using the code of Zia & Lecampion (2020) and summarised in Table 2. These comprise three parametric cases, each with three different injection volumes and rates. The plots stop at the point when the simulated fracture reached the edge of the meshed domain or Pyfrac terminated the simulation.

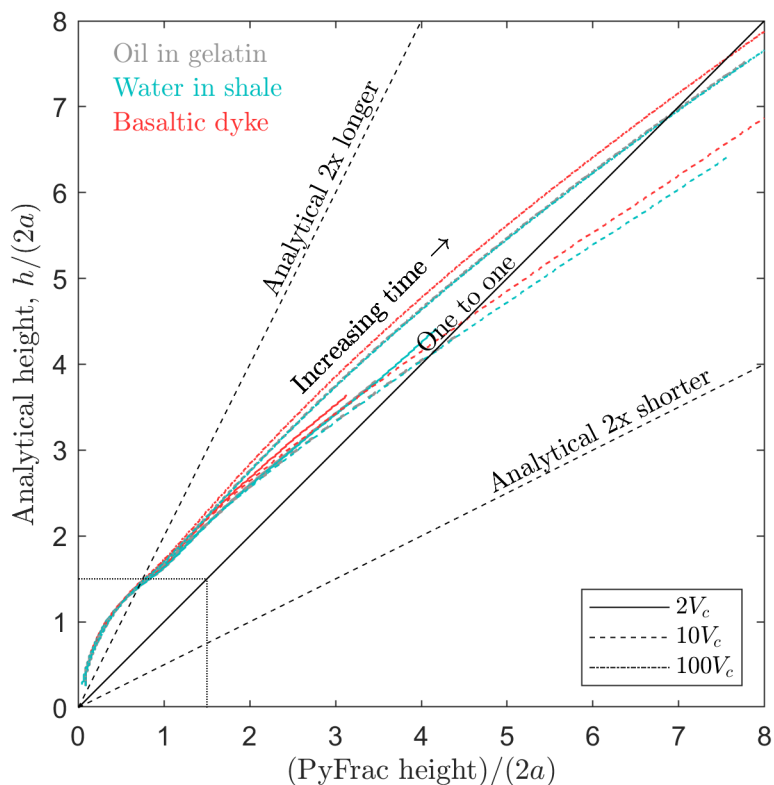


Figure 4: Numerical versus analytical (3.5) height estimates at times  $t$  since injection. Height here is defined as the distance from the injection point to the upper tip. This plot uses the same numerical solutions as in Fig. 3.

Quantity	Symbol	Unit	Exp. 1933	Exp. 1945	Exp. 1967
Young's modulus	$E$	Pa	1995	306	595
Poisson's ratio	$\nu$		0.5	0.5	0.5
Fracture toughness	$K_C$	$\text{Pa m}^{1/2}$		17	23.8
Oil/gelatin density difference	$\Delta\gamma$	$\text{kg m}^{-3}$	260	160	150
Oil viscosity	$\eta$	Pa s	$1.74 \times 10^{-3}$	$48 \times 10^{-3}$	$970 \times 10^{-3}$
Injected Volume	$V_I$	$\text{m}^3$	40	10	10
Critical Volume (2.2)	$V_c$	$\text{m}^3$		$2.64 \times 10^{-5}$	$3.72 \times 10^{-5}$
Predicted velocity (3.4)	$\bar{v}_V$	$\text{m s}^{-1}$	1.69	0.038	$8.4 \times 10^{-4}$
Reynolds number (3.8)	Re		$3.1 \times 10^3$	2.81	0.0022

Table 3: Properties of the analog gelatin experiments of Smittarello (2019). Experiment reference numbers are shown in the first row. Note that the injection rates were not recorded.

#### 4.2. Comparison to analog oil injection gelatin experiments

410

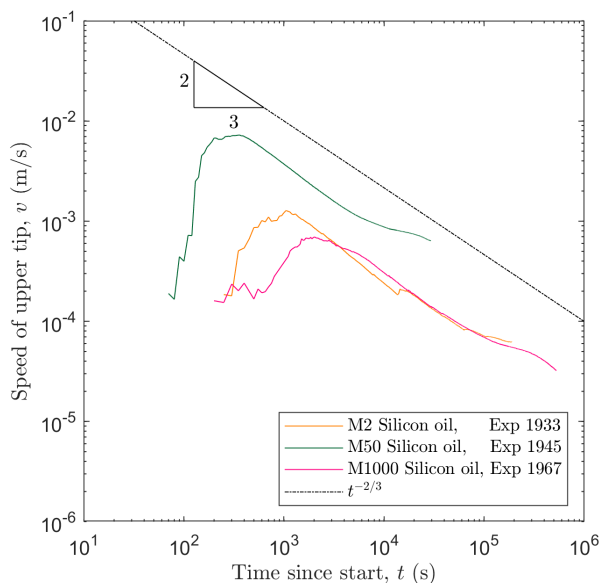
411 The gelatin experiments of Heimpel & Olson (1994) and Smittarello (2019) show that crack  
 412 ascent rates are related to the injected volume. Our analysis can be used to understand  
 413 what determines the speed of fractures in such experiments. We reiterate that our predicted  
 414 maximum ascent speed (3.4) is an overestimate for cracks filled with low-viscosity fluids,  
 415 such as air, that are prone to turbulence (as is the case in many analog experiments). The  
 416 Reynolds number in (3.8) can be used to assess whether the flow in the fracture is turbulent,  
 417 which we have shown is generally true for gelatin experiments with air-filled fractures.  
 418 Additionally, many fluids typically injected such as air and water, are non-wetting fluids when  
 419 in contact with gelatin solids. When these flow through fractures in gelatin, the lubrication  
 420 approximation breaks down; all fluid can be expelled from between the walls. Both effective  
 421 closure of the fracture's tail and near-constant ascent rates are indications that non-wetting  
 422 processes are active. This has led some authors to conclude that typical experimental fluids  
 423 are not a good analog for the fluids that drive fracture in the crust (Taisne & Tait 2009).

424 Silicon oils are wetting with respect to the gelatin. In Fig. 5 we plot the speed of silicon  
 425 oils hydro-fracturing upward through gelatin solids. The experiments are described in  
 426 Smittarello (2019); the material properties are given in Table 3. These experiments have  
 427 well-constrained injection volumes and elastic parameters, and therefore provide a suitable  
 428 basis for comparison with our equations.

429 Figure 5a shows that after the post-injection transient ends, the ascent speed of the  
 430 experimental fractures follow the  $t^{-2/3}$  trend. Fig. 5b shows that our analysis predicts the rate  
 431 of upward propagation in the buoyancy-driven regime within a factor of two of the observed  
 432 speed. This good match appears to be independent of the rate of injection, presence of tank  
 433 walls and free surface, loads placed on the gelatin during the experiments and the evolving  
 434 elastic properties during the time span of the experiments as the gelatin warms (Smittarello  
 435 2019). Note that the fracture accelerates as it approaches the free surface. In Fig. 5a, this  
 436 appears as a deviation from  $t^{-2/3}$  trend toward reduced deceleration at later times (Rivalta &  
 437 Dahm 2006). The line representing experiment 1933, where the prediction is off by an order  
 438 of magnitude, shows the effect of a high Reynolds number (Table 3), which significantly  
 439 reduces the ascent speed of the fracture relative to that predicted by laminar flow.



a)



b)

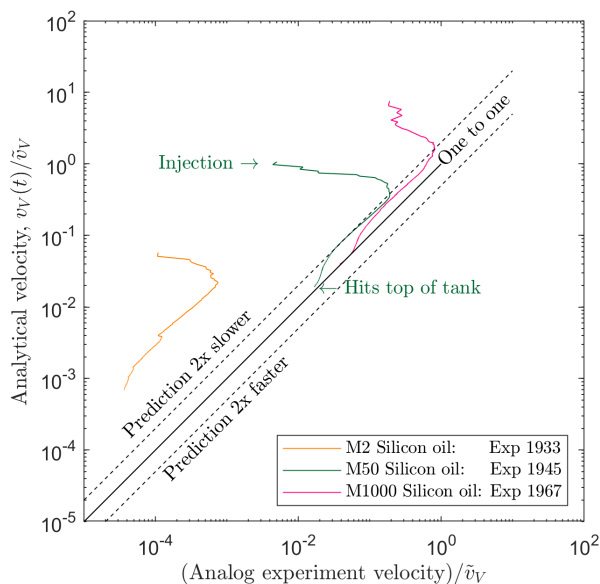


Figure 5: Upper tip ascent speed from analog experiments where silicon oil is injected into in gelatin solids. Parameter values are given in table 3. a) Ascent rate versus time, showing that at late times, curves approach the  $t^{-2/3}$  asymptote (dash-dot). b) The speed from the experiments is plotted against the predicted speed using equation (3.6), where the time is that elapsed since the start of the injection.

## 440 5. Discussion

### 441 5.1. Some insights into ascent velocity

442 In our numerical simulations, we have made sure that the injection rates are such that the  
443 specified fluid volume has been injected into the fracture by the time the crack height reaches  
444  $2a$ . This means that at the end of the injection, the crack tip-line is approximately circular  
445 and  $\bar{v}_V$  should approximate the ascent velocity at this time. We expect that if the time scale of  
446 injection is much lower than the time scale of propagation, that once the crack has exceeded  
447 the critical radius of Davis *et al.* (2020), it will begin to elongate in the direction of the stress  
448 gradient. In these cases, the upwards ascent speed should lie somewhere between the two  
449 limiting regimes; that predicted by equation (2.5) of a finite fluid batch, and the ascent rate  
450 defined by a constant fluid flux, equation 2.4 (Möri & Lecampion 2021*b*). Even when the  
451 injection rate is low, once the entire fluid batch has been injected into the fracture and as  
452  $t$  increases, our analysis to predict the deceleration rate will remain valid. One example of  
453 this can be seen in the analog gelatin data in Fig. 5. Here, for M50 silicon oil the ascent rate  
454 doesn't reach  $\bar{v}_V$  because the injection rate was low. Despite this, our equation predicts the  
455 decay in the speed at later times.

456 Mutch *et al.* (2019) use geochemical techniques to retrieve magma ascent rates of the  
457 Borgarhraun eruption, northern Iceland. These results suggest the magma's ascent through  
458 the crust was rapid, in the range of 0.02 to 0.1 m s<sup>-1</sup>. We now test to see if our equations can  
459 correctly predict such ascent speeds. The erupted lava volume of Borgarhraun was reported  
460 between 0.014–0.14 km<sup>3</sup> and the magma density was around 2700 kg m<sup>-3</sup> (Maclennan *et al.*  
461 2003; Hartley & Maclennan 2018). Assuming the following parameters:  $\rho_r = 2750$  kg m<sup>-3</sup>,  
462  $E = 10$ –40 GPa,  $\nu = 0.25$  and  $\eta = 10$ –30 Pa s, we find the maximum ascent rate from equation  
463 (3.4) between 0.08 and 9.4 m s<sup>-1</sup>. Calculating the average speed from equation (3.6) between  
464  $2a$  and the reported distance traversed by the batch of 24 km, we find that it is between 0.06  
465 and 9.6 m s<sup>-1</sup>. Hence we observe that by using approximate crustal parameters, our analysis  
466 provides a simple means to predict and explain how a relatively small batch of magma can  
467 traverse the crust within a week.

468 Our results are also important in the context of hydro-fracturing operations. They can be  
469 used to quantify the time it would take for a fluid to pass into overlying formations and,  
470 furthermore, they give an estimate of the area of rock exposed to the crack surfaces. We  
471 aim here to give an indication of how this formula can be applied to industrial operations  
472 such as hydro-fracturing. We envision the case of injecting a fluid volume of 25 m<sup>3</sup>, where  
473 the fluid viscosity ranges between 10<sup>-3</sup>–10<sup>-2</sup> Pa s, the rock stiffness's is 10–40 GPa and  
474 assuming the rock and fluid density are 2700 kg m<sup>-3</sup> and 1000 kg m<sup>-3</sup>, respectively. For  
475 this range of properties, using equation (3.5), it would take between 15 minutes to 5 hours  
476 for the fracture to propagate 1000 m vertically and 2 to 40 hours to propagate 2000 m,  
477 noting that for the faster ascent rates, turbulent flow may occur in the early stages of ascent.  
478 These ascent rates and distances suggest this is an efficient way to transport fluid through  
479 the crust, and they beg the question, which processes act to slow or stop this ascent? Low-  
480 viscosity fluids in porous formations can leak into the surrounding rock. This process is  
481 known as fluid leak-off, this would reduce the effective volume driving the fracture upwards  
482 and could, in some cases, change the dynamics of flow-driven fracture growth within the tip  
483 (Dontsov & Peirce 2015). Note, that in the context of a dyke, solidification of the magma  
484 along the dyke walls is mathematically equivalent to leak-off. This process is well understood  
485 for constant inflow 2D buoyant cracks (Lister 1994*b*). It is of interest to quantify how the  
486 upward propagation speed and trapping height for a finite volume of fluid is changed by fluid  
487 leak-off/solidification during ascent (Detournay 2004). Here we leave this extended analysis  
488 to future studies where more experimental data is available, but note that Lister (1994*a*,

489 Eq.5.1) provides a 2D approximation of the maximal distance a fixed finite-area dyke can  
490 propagate when solidification is considered. Recall that in Sec. 3.2 we have shown how to  
491 convert similar equations to 3D.

## 492 6. Conclusions

493 We have provided an analytical approximation of the maximum ascent speed of a three-  
494 dimensional, buoyant fluid-driven fracture containing a finite fluid volume. We verified this  
495 by comparison with outputs from a hydro-fracture simulator. We showed that the ascent speed  
496 decays away from this maximum due to viscous drag in the growing tail region, at a rate  
497 that is asymptotic to  $t^{-2/3}$ . Our quantitative approximations help to explain why a dyke can  
498 traverse the crust in a time that is of order days. They also bring hydro-fracturing operations  
499 into question, suggesting typical injection volumes should ascend through the crust unless  
500 another process acts to trap the fracture or drain off fluid.

## 501 7. Acknowledgements

502 This research received funding from the European Research Council under Horizon 2020  
503 research and innovation program grant agreement number 772255 and the DFG-ICDP grant  
504 N. RI 2782/3-1. We thank Francesco Maccaferri and Virginie Pinel for helping both in the  
505 design and execution the analog crack ascent experiments in gelatin. We are also grateful  
506 to Saeed Salimzadeh who provided us with data from his numerical experiments and Brice  
507 Lecampion and his group on their responses to questions about PyFrac.

## 508 8. Declaration of Interests

509 The authors report no conflict of interest.

## 510 9. Data and codes

511 The codes use to produce the figures and run the PyFrac simulations are at: <https://doi.org/10.5281/zenodo.6669935>. The data and movies of the gelatin experiments and data  
512 from the PyFrac simulations are at: <https://doi.org/10.5281/zenodo.6669974>.

## 514 Appendix A. 2D similarity solutions - fixed finite-area

515 We restate the solutions of Roper & Lister (2007) in dimensional coordinates. The half height  
516 of a Weertman crack is (Pollard & Muller 1976)

$$517 \quad c_h = \left( \frac{K_c}{\Delta\gamma\sqrt{\pi}} \right)^{2/3}. \quad (\text{A } 1)$$

518 The head length is  $2c_h$  and it's areal is

$$519 \quad A_h = \frac{K_c^2(1-\nu)}{2\mu\Delta\gamma}. \quad (\text{A } 2)$$

520 If the initial crack area is  $A$ , then the tail area is  $A_t = A - A_h$ . Roper & Lister (2007) define  
521 the fracture's entire tail length as

$$522 \quad z_t^+ = \left( \frac{9A_t^2\Delta\gamma t}{16\eta} \right)^{1/3}. \quad (\text{A } 3)$$

523 The opening profile (wall separation) of the head ( $w_h$ ) is

$$524 \quad \frac{w_h}{2} = \frac{(1-\nu)K_c}{2\mu} \sqrt{\frac{c_h}{\pi}} \sqrt{1 - \left(\frac{z_h}{c_h}\right)^2} \left(1 + \frac{z_h}{c_h}\right), \quad (\text{A } 4)$$

525 where  $z_h$  spans from  $-c_h$  to  $c_h$ . The separation of the walls in the tail ( $w_t$ ) is

$$526 \quad \frac{w_t}{2} = \sqrt{\frac{\eta z_t}{\Delta\gamma t}}, \quad (\text{A } 5)$$

527 where here  $z_t$  spans from 0 to  $z_t^+$  (from the base of the crack to where this meets the head  
528 solution). The two solutions are then joined by moving  $z_h$  upwards and neglecting parts  
529 thinner than the top of the tail at the base of the head solution. We remind readers that here,  
530 the crack's centre at time  $t = 0$  lies at  $z = c_h$  (the injection point).

#### REFERENCES

- 531 CARTWRIGHT, J., KIRKHAM, C., FOSCHI, M., H, NEIL., RODRIGUEZ, K. & JAMES, D. 2021 Quantitative  
532 reconstruction of pore-pressure history in sedimentary basins using fluid escape pipes. *Geology*  
533 **49** (5), 576–580.
- 534 DAHM, T. 2000 On the shape and velocity of fluid-filled fractures in the earth. *Geophysical Journal*  
535 *International* **142** (1), 181–192.
- 536 DAS, S.B., JOUGHIN, I., BEHN, M.D., HOWAT, I.M., KING, M.A., LIZARRALDE, D. & BHATIA, M. 2008  
537 Fracture propagation to the base of the greenland ice sheet during supraglacial lake drainage. *Science*  
538 **320** (5877), 778–781.
- 539 DAVIS, T., RIVALTA, E. & DAHM, T. 2020 Critical fluid injection volumes for uncontrolled fracture ascent.  
540 *Geophysical Research Letters* **47** (14), e2020GL087774.
- 541 DETOURNAY, E. 2004 Propagation regimes of fluid-driven fractures in impermeable rocks. *International*  
542 *Journal of Geomechanics* **4** (1), 35–45.
- 543 DONTSOV, E.V. & PEIRCE, A.P. 2015 A non-singular integral equation formulation to analyse multiscale  
544 behaviour in semi-infinite hydraulic fractures. *Journal of Fluid Mechanics* **781**.
- 545 EPA, US 2016 Hydraulic fracturing for oil and gas: Impacts from the hydraulic fracturing water cycle on  
546 drinking water resources in the united states. *Washington, DC: US Environmental Protection Agency,*  
547 *EPA/600/R-16/236F* .
- 548 GERMANOVICH, L.N., GARAGASH, D., MURDOCH, L.C. & ROBINOWITZ, M. 2014 Gravity-driven hydraulic  
549 fractures. In *AGU Fall Meeting Abstracts*, , vol. 2014, pp. H53C–0874.
- 550 HARTLEY, M. & MACLENNAN, J. 2018 Magmatic densities control erupted volumes in icelandic volcanic  
551 systems. *Frontiers in Earth Science* **6**, 29.
- 552 HEIMPEL, M. & OLSON, P. 1994 Buoyancy-driven fracture and magma transport through the lithosphere:  
553 models and experiments. In *International Geophysics*, , vol. 57, chap. 10, pp. 223–240. Elsevier.
- 554 LISTER, J.R. 1990 Buoyancy-driven fluid fracture: the effects of material toughness and of low-viscosity  
555 precursors. *Journal of Fluid Mechanics* **210**, 263–280.
- 556 LISTER, J.R. 1994a The solidification of buoyancy-driven flow in a flexible-walled channel. part 1. constant-  
557 volume release. *Journal of Fluid Mechanics* **272**, 21–44.
- 558 LISTER, J.R. 1994b The solidification of buoyancy-driven flow in a flexible-walled channel. part 2. continual  
559 release. *Journal of Fluid Mechanics* **272**, 45–66.
- 560 LISTER, J.R. & KERR, R.C. 1991 Fluid-mechanical models of crack propagation and their application to  
561 magma transport in dykes. *Journal of Geophysical Research: Solid Earth* **96** (B6), 10049–10077.
- 562 MACLENNAN, J., MCKENZIE, D., HILTON, F., GRONVÖLD, K. & SHIMIZU, N. 2003 Geochemical variability in  
563 a single flow from northern iceland. *Journal of Geophysical Research: Solid Earth* **108** (B1), ECV–4.
- 564 MAIR, R., BICKLE, M., GOODMAN, D., KOPPELMAN, B., ROBERTS, J., SELLEY, R., SHIPTON, Z., THOMAS, H.,  
565 WALKER, A., WOODS, E. & YOUNGER, P. 2012 Shale gas extraction in the uk: a review of hydraulic  
566 fracturing. *Tech. Rep.*. Royal Society and Royal Academy of Engineering.
- 567 MÖRI, A & LECAMPION, B 2021a Arrest of a radial hydraulic fracture upon shut-in of the injection.  
568 *International Journal of Solids and Structures* **219**, 151–165.

- 569 MÖRI, A. & LECAMPION, B. 2021*b* Limiting regimes of a three-dimensional buoyant hydraulic fracture  
570 emerging from a point source. In *AGU Fall Meeting 2021*.
- 571 MÖRI, A., LECAMPION, B. & ZIA, H. 2020 Fully-coupled 3d modelling of magmatic dike propagation-finite  
572 pulse release from a point source. *Tech. Rep.*. Copernicus Meetings.
- 573 MOUKHTARI, F.E., LECAMPION, B. & ZIA, H. 2020 Planar hydraulic fracture growth perpendicular to the  
574 isotropy plane in a transversely isotropic material. *Journal of the Mechanics and Physics of Solids*  
575 **137**, 103878.
- 576 MUTCH, E.J.F., MACLENNAN, J., SHORTTLE, O., EDMONDS, M. & RUDGE, J.F. 2019 Rapid transcrustal magma  
577 movement under iceland. *Nature Geoscience* **12** (7), 569–574.
- 578 OKAMOTO, A. & TSUCHIYA, N. 2009 Velocity of vertical fluid ascent within vein-forming fractures. *Geology*  
579 **37** (6), 563–566.
- 580 PEIRCE, A. 2015 Modeling multi-scale processes in hydraulic fracture propagation using the implicit level  
581 set algorithm. *Computer Methods in Applied Mechanics and Engineering* **283**, 881–908.
- 582 PEIRCE, A. 2016 Implicit level set algorithms for modelling hydraulic fracture propagation. *Philosophical*  
583 *Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences* **374** (2078),  
584 20150423.
- 585 POLLARD, D.D. & MULLER, O.H. 1976 The effect of gradients in regional stress and magma pressure on the  
586 form of sheet intrusions in cross section. *Journal of Geophysical Research* **81** (5), 975–984.
- 587 POLLARD, D.D. & TOWNSEND, M.R. 2018 Fluid-filled fractures in earth's lithosphere: Gravitational loading,  
588 interpenetration, and stable height of dikes and veins. *Journal of Structural Geology* **109**, 38–54.
- 589 RIVALTA, E. & DAHM, T. 2006 Acceleration of buoyancy-driven fractures and magmatic dikes beneath the  
590 free surface. *Geophysical Journal International* **166** (3), 1424–1439.
- 591 RIVALTA, E., TAISNE, B., BUNGER, A.P. & KATZ, R.F. 2015 A review of mechanical models of dike  
592 propagation: Schools of thought, results and future directions. *Tectonophysics* **638**, 1–42.
- 593 ROPER, S.M. & LISTER, J.R. 2007 Buoyancy-driven crack propagation: the limit of large fracture toughness.  
594 *Journal of Fluid Mechanics* **580**, 359–380.
- 595 RUBIN, A.M. 1995 Propagation of magma-filled cracks. *Annual Review of Earth and Planetary Sciences*  
596 **23** (1), 287–336.
- 597 SALIMZADEH, S., ZIMMERMAN, R.W. & KHALILI, N. 2020 Gravity hydraulic fracturing: A method to create  
598 self-driven fractures. *Geophysical Research Letters* p. e2020GL087563.
- 599 SCHULTZ, R.A. 2016 Causes and mitigation strategies of surface hydrocarbon leaks at heavy-oil fields:  
600 examples from alberta and california. *Petroleum Geoscience* **23** (2), 231–237.
- 601 SECOR, D.T. & POLLARD, D.D. 1975 On the stability of open hydraulic fractures in the earth's crust.  
602 *Geophysical Research Letters* **2** (11), 510–513.
- 603 SMITTARELLO, D. 2019 Propagation des intrusions basaltiques. PhD thesis, Université Grenoble Alpes.
- 604 SMITTARELLO, D., PINEL, V., MACCAFERRI, F., FURST, S., RIVALTA, E. & CAYOL, V. 2021 Characterizing the  
605 physical properties of gelatin, a classic analog for the brittle elastic crust, insight from numerical  
606 modeling. *Tectonophysics* **812**, 228901.
- 607 SPENCE, D.A. & TURCOTTE, D.L. 1990 Buoyancy-driven magma fracture: A mechanism for ascent through  
608 the lithosphere and the emplacement of diamonds. *Journal of Geophysical Research: Solid Earth*  
609 **95** (B4), 5133–5139.
- 610 STEVENS, L.A., BEHN, M.D., MCGUIRE, J.J., DAS, S.B., JOUGHIN, I., HERRING, T., SHEAN, D.E. & KING,  
611 M.A. 2015 Greenland supraglacial lake drainages triggered by hydrologically induced basal slip.  
612 *Nature* **522** (7554), 73–76.
- 613 TADA, H., PARIS, P. & IRWIN, G.R. 2000 *The Stress Analysis of Cracks Handbook; Third edition*, , vol. 130.  
614 New York: ASME Press.
- 615 TAISNE, B. & TAIT, S. 2009 Eruption versus intrusion? arrest of propagation of constant volume, buoyant,  
616 liquid-filled cracks in an elastic, brittle host. *Journal of Geophysical Research: Solid Earth* **114** (B6).
- 617 TOLSTOY, M., COWEN, J.P., BAKER, E.T., FORNARI, D.J., RUBIN, K.H., SHANK, T.M., WALDHAUSER, F.,  
618 BOHNENSTIEHL, D.R., FORSYTH, D.W., HOLMES, R.C. & LOVE, B. 2006 A sea-floor spreading event  
619 captured by seismometers. *Science* **314** (5807), 1920–1922.
- 620 WEERTMAN, J. 1971 Theory of water-filled crevasses in glaciers applied to vertical magma transport beneath  
621 oceanic ridges. *Journal of Geophysical Research* **76** (5), 1171–1183.
- 622 ZIA, H. & LECAMPION, B. 2020 Pyfrac: A planar 3d hydraulic fracture simulator. *Computer Physics*  
623 *Communications* **255** (107368), 1–11.
- 624 ZIA, H., LECAMPION, B. & ZHANG, W. 2018 Impact of the anisotropy of fracture toughness on the propagation  
625 of planar 3d hydraulic fracture. *International Journal of Fracture* **211** (1), 103–123.