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1 **Title:** Heterogenous controls on lake color and trends across the high-elevation U.S. Rocky Mountain
2 region

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14 **Running head:** Lake color trends Rockies

15 **Keywords:** Landsat; trend analysis; long-term trends; oligotrophic; mountain lakes; water quality,
16 climate change

17 **Abstract**

18 Global change may contribute to ecological changes in high-elevation lakes and reservoirs,
19 but a lack of data makes it difficult to evaluate spatiotemporal patterns. Remote sensing imagery can
20 provide more complete records to evaluate whether consistent changes across a broad geographic
21 region are occurring. We used Landsat surface reflectance data to evaluate spatial patterns of
22 contemporary lake color (2010-2020) in 940 lakes in the U.S. Rocky Mountains. Intuitively, we found
23 that most of the lakes in the region are blue (66%) and were found in steep-sided watersheds (>22.5°)

24 or alternatively were relatively deep (>4.5m) with mean annual air temperature (MAAT) <4.5°C. Most
25 green/brown lakes were found in relatively shallow sloped watersheds with MAAT ≥4.5°C. We
26 extended the analysis of contemporary lake color to evaluate changes in color from 1984-2020 for a
27 subset of lakes with the most complete time series (n=527). We found limited evidence of lakes shifting
28 from blue to green states, but rather, 55% of the lakes had no trend in lake color, 32% of lakes were
29 trending toward bluer wavelengths, and only 13% shifted toward greener wavelengths. Lakes and
30 reservoirs with the most substantial shifts toward blue wavelengths tended to be in urbanized, human
31 population centers at relatively lower elevations. In contrast, lakes that shifted to greener wavelengths
32 did not relate clearly to any lake or landscape features that we evaluated, though declining winter
33 precipitation and warming summer and fall temperatures may play a role in some systems.
34 Collectively, these results suggest that the interactions between local landscape factors and broader
35 climatic changes can result in heterogeneous, context-dependent changes in lake color.

36 **Introduction**

37 High-elevation lakes and reservoirs form the basis of a critical water supply network for arid
38 and semi-arid cities and communities downstream. However, climate change threatens these
39 ecosystems via altered temperature and precipitation regimes (Christianson et al., 2020; Maberly et
40 al., 2020), lake ice phenology (Benson et al., 2012; Preston et al., 2016), lake temperature
41 (Christianson et al., 2019; Sadro et al., 2018; Smits et al., 2020) and, in turn, ecosystem function and
42 biological composition. In addition to climate change, increasing nutrient loading presents an additional
43 steady change that can lead to increased algal production (Moser et al., 2019; Oleksy et al., 2021).

44 Despite these potential threats to high-elevation lakes, examining shifts in freshwater
45 ecosystems at large spatial scales is challenging because of inadequate coverage and a strong bias
46 of analyses towards a few well-monitored lakes (Stanley et al., 2019). Physiochemical changes (e.g.,
47 ice-cover duration, water chemistry, surface temperature) in a number of pristine high-elevation lakes
48 suggests that these shifts are significant (Moser et al., 2019; Preston et al., 2016). Summer warming

49 in combination with nitrogen deposition is leading to algal assemblage shifts and increasing
50 productivity in lakes along the Colorado Front Range (Oleksy et al., 2020). In addition, snowpack and
51 summer weather conditions are strong controls on water chemistry and algal biomass for mountain
52 lakes (Oleksy et al., 2020; Preston et al., 2016; Sadro et al., 2018). While there has been recent
53 research examining regional to continental scale changes in lake nutrients (Oliver et al., 2017;
54 Stoddard et al., 2016), water clarity (Topp et al., 2021), lake color (Kuhn and Butman, 2021), and algal
55 blooms (Wilkinson et al., 2021) there have been no regional studies, to our knowledge, on high-
56 elevation lake shifts likely due to a lack of *in situ* water quality monitoring data (Read et al., 2017).

57 While remote sensing can be used to directly estimate water quality parameters (Topp et al.,
58 2020), lake water color is relatively easy to infer from satellite and is less prone to prediction errors
59 (Giardino et al., 2014). Color is also an intuitive and integrative metric that can serve as an indicator
60 of water quality parameters, including colored dissolved organic matter, which can be used to infer
61 estimates of total organic carbon, dissolved organic carbon (Ouyang et al., 2006), chlorophyll-a (proxy
62 for algal productivity; Cao et al., 2020), and suspended sediment (Dekker et al., 2001).

63 Here we used satellite-derived lake color to address three core objectives to better understand
64 lake color in the U.S. Rocky Mountains:

- 65 1. We evaluated the contemporary spatial distribution of average summer lake color.
- 66 2. We quantified how lake color has changed in the region since the beginning of the Landsat
67 record (1984).
- 68 3. We examined which lake, landscape, and climatological features of lakes relate to spatial
69 patterns and temporal trends in lake color.

70 Through these three objectives we aimed to understand the current patterns of lake color across the
71 U.S. Rocky Mountain region and to assess how climate change and other disturbances might be
72 changing or will change color, and therefore, lake ecology and related ecosystem characteristics.

73 **Methods**

74 *Lake color*

75 We used remote sensing data from the LimnoSat-US database (Topp et al., 2020), a robust
76 collection of Landsat surface reflectance data for 56,792 U.S. lakes. The LimnoSat-US data extracts
77 USGS Tier 1 surface reflectance values over Landsat 5, Landsat 7, and Landsat 8 sensors dating
78 back to 1984 from the deepest point of lakes, or the point furthest from any shoreline. All the Landsat
79 imagery has been atmospherically corrected, and then adjusted so each satellite had unbiased data
80 across time and between satellites (Topp et al., 2020). We limited the analysis to high elevation lakes
81 in the Rocky Mountain Region, which we define as the parts of Idaho, Colorado, Montana, Wyoming,
82 Utah, and New Mexico above 1400m. This captures many mountain lakes in the region as well as
83 high-elevation plains lakes and reservoirs. We examined spatiotemporal patterns in the color of the
84 lake as perceived by the human eye, called the Dominant Wavelength (DWL), which maps directly to
85 the Forel-Ule scale, a water transparency classification scale (Wernand and Van der Woerd, 2010).
86 For both color-measuring approaches, blue lakes (<520nm) are generally considered oligotrophic,
87 while change in color from blue to green wavelengths generally corresponds to shifts in trophic state
88 from mesotrophic to eutrophic (>520nm). Color changes from green toward brown wavelengths can
89 indicate either a dystrophic system or a eutrophic lake with high suspended sediment in the water
90 column (>575nm).

91 *Classification of spatial patterns*

92 To understand broad-scale spatial patterns, we examined the median contemporary (2010-
93 2020) lake color across the U.S. Rocky Mountains. We included data from the summer period (July 1-
94 September 15) to minimize seasonal variation and the impact of snow and ice cover, which can persist
95 into June for some of the highest elevation lakes. We joined the LimnoSat-US lake color data to the
96 National Hydrography Dataset (U.S. Geologic Survey, 2021), the Global Lake Area, Climate, and
97 Population dataset (Smith et al., 2021), watershed-level metrics from the LakeCat database (Hill et al.,

98 2018), LAGOS-US NETWORKS (King et al., 2021) and LAGOS-US Reservoir (Polus et al., 2021). We
99 used information about the lake, landscape, lake type (natural lake or reservoir), and connectivity
100 features from these datasets to explain lake color spatial patterns in lakes that spanned a broad range
101 of environmental contexts (Table 1).

102 We divided the population of lakes into two categories: blue/clear (DWL <530 nm; n=620) or
103 green/turbid (DWL >530 nm; n=320; Figure 2A). To address our first research objective, we used a
104 Classification and Regression Tree (CART; Therneau and Atkinson, 1997) to determine which
105 environmental characteristics explained variation in lake color across the region using the *rpart*
106 package (Therneau and Atkinson, 2019). The training dataset included 80% of the total population
107 (n=765). We calculated the global model accuracy by predicting lake color groupings for the out-of-
108 sample lakes and assessed model performance with a confusion matrix. We visualized the results
109 with the *cvms* (Olsen and Zachariae, 2021) and *ggparty* (Borkovec and Madin, 2019) R packages for
110 the confusion matrix and decision trees, respectively. All analyses and data visualizations were done
111 in R version 4.0.5 (R Core Team, 2021).

112 *Trend analysis*

113 For the trend analysis, we built a separate dataset that included only lakes that had at least 3
114 cloud-free summer images for a minimum of 30 consecutive years between 1984-2020 for a total of
115 527 lakes in the analysis. This accounts for approximately a quarter of all lakes in this region that are
116 greater than 10 ha in area and over 1400 m in elevation (Figure S1). We calculated the non-parametric
117 Theil-Sen's slope for each lake time series of median summer color using the *trend* package (Pohlert,
118 2020). We used the Mann-Kendall z-score and compared the p-value from that z-score to $\alpha = 0.05$.
119 We categorized each lake into one of five possible trend categories:

120 **1) No trend** when the p-value of the Sen's slope was greater than 0.05. All other categories
121 had p-values of <0.05;

122 **2) Blue->Greener**, for lakes that started blue during the first half of the record (median DWL

123 <530 nm; 1984-2005) and had a positive slope;
124 **3) Intensifying Green/brown** for lakes that started green prior to 2005 (median DWL >530
125 nm) and had a positive slope;
126 4) **Green->Bluer** for lakes that started green (median DWL >530 nm between 1984-2005)
127 and had a negative slope; and
128 **5) Intensifying Blue** for lakes that started blue prior to 2005 (DWL <530 nm) and had a
129 negative slope.

130 For lakes in the Blue->Greener and Green->Bluer categories, we assessed whether the median lake
131 color in the later part of the record indicated a modal shift in color from predominantly blue to
132 predominantly green/brown, or vice versa, consistent with the spatial color categorization.

133 We conducted a Random Forest analysis to explore the drivers of color trends (Breiman,
134 2001). Here, we grouped together all lakes with positive trends (Intensifying Green/Yellow and Blue-
135 >Greener) and negative trends (Intensifying Blue and Green->Bluer) into composite categories for a
136 total of three trend categories (Negative, No Trend, Positive). Predictors included all those considered
137 in the spatial CART described above (Table 1) as well as changes in seasonal precipitation,
138 temperature, and human population size. We used the *prism* package to download the daily estimate
139 of temperature and precipitation from the Oregon Parameter-elevation Relationships on Independent
140 Slopes Model (PRISM) project (Hart and Bell, 2015). For each lake-year, we calculated the mean
141 winter (December-February), spring (March-May), summer (June-August), and fall (September-
142 November) temperature and precipitation. Then, we calculated the Sen's slope of temperature and
143 precipitation for each lake and season from 1984-2020.

144 We built the random forest models using the `rand_forest` function in the *parsnip* package using
145 the "ranger" engine (Kuhn and Vaughan, 2021a; Wright and Ziegler, 2017). We randomly chose 60%
146 of the data as our training data set and 40% as our test dataset which ensured that at least 25% of the
147 observations in each trend category were set aside for validation. We tuned the two hyperparameters

148 using ten-fold cross-validation. The optimum number of predictors at each node ($mtry = 4$) and the
149 minimum n to split at any node ($min_n = 3$) for the final model was selected according to the best
150 Receiver Operating Characteristic curve and overall classification accuracy using the *yardstick*
151 package (Kuhn and Vaughan, 2021b). The final random forest model consisted of 1000 trees and was
152 evaluated on the validation data. We present the top 10 predictors based on Variable Importance (VI),
153 computer as the total decrease in node impurity averaged over all trees.

154 **Results**

155 *Spatial patterns*

156 Our dataset included 940 lakes above 1400m across the six-state Rocky Mountain region (Figure 1).
157 Between 2010-2020, 66% of the lakes were predominantly blue ($n=620$) while 34% of the lakes were
158 predominantly green/brown ($n=320$; Figure S2). The CART analysis revealed that watershed slope,
159 mean annual air temperature (MAAT), and maximum lake depth were important determinants of lake
160 color (Figure 2C). Most green/brown lakes were found in relatively shallow sloped watersheds with
161 MAAT $\geq 4.5^\circ\text{C}$. Lakes situated in watersheds with slope angles $\geq 22.5^\circ$ were most likely to be classified
162 as blue lakes. Similarly, another set of blue lakes were common in less steep watersheds with MAAT
163 $\leq 4.5^\circ\text{C}$ with maximum depth ≥ 4.5 meters while shallower lakes ($< 4.5\text{m}$ deep) in those areas were
164 more likely to be green/brown. Overall, the CART model was able to correctly classify 84% of blue
165 lakes and 68% of green/brown lakes in the test dataset (Figure 2B).

166 *Cross-lake color trends*

167 In the U.S. Rocky Mountains, we detected no trends in lake color between 1984-2020 in 55% of lakes
168 ($n=290$, Figure 3). However, 32% of lakes were trending bluer ($n=166$) and reservoirs showed the
169 largest improvements in water quality. Specifically, 71% of the lakes that trended bluer were reservoirs
170 ($n=30$), and 75% of the lakes that were intensifying blue were reservoirs ($n=30$). Most of the lakes
171 trending from Green->Bluer were in Colorado (71% or $n=72$; Figure 4), including many Colorado

172 reservoirs that switched from Green/brown to blue/clear (n=14, Table S1). Median lake color shifted
173 toward greener wavelengths in 13% of the population of lakes (n=71), with 34 lakes in the Intensifying
174 Green/brown category and 37 lakes in the Blue->Greener category. Of the Blue->Greener lakes, six
175 of them crossed the 530nm threshold consistently in recent years such that they were classified as
176 Green/brown in the spatial analysis.

177 Although our Random Forest model poorly predicted lake greening (Figure S3), a combination
178 of static variables and climatic trends partially explained some trends in lake color (Figure 5). The
179 variables with highest importance included total human population in the lake-watershed (Variable
180 Importance=6.07), lake elevation (VI=5.91), changes in winter precipitation (VI=5.58), urban landcover
181 (VI=4.4), and changes of spring temperature (VI=4.25). The majority of the lakes that were Intensifying
182 Blue or trending Green->Bluer were located in relatively urbanized watersheds with some of the
183 highest human population densities in the region (Figure 5A,D). These lakes also tended to be located
184 at lower elevations relative to lakes not experiencing color shifts or lakes that were greening (Figure
185 5B). Both greening and blueing lakes were associated with decreases in winter precipitation between
186 1984-2020 compared to lakes with no trend (Figure 5C). Furthermore, blueing lakes tended to be in
187 areas where spring air temperatures were cooling slightly relative to greening lakes or lakes without
188 color changes (Figure 5E), though notably for both the climatic variables only a small subset of the
189 trends were statistically significant (Figure S4, Table S2).

190 Overall, the most widespread climatic trends in the region were increasing summer and fall
191 temperatures (Table S2). Although increasing fall temperatures were widespread in this region, there
192 were no differences among color trend groups (ANOVA $F_{2,328}=2.55$, $p=0.08$). However, absolute rates
193 of summer warming varied among color groups (Kruskal–Wallis H-test, $p<0.001$). Specifically, since
194 1984, summer temperatures increased on average 0.23°C more in lakes with no change in color
195 compared to lakes that were trending blue (95% CI: 0.06-0.4°C). Further, rates of summer warming
196 were 0.34°C higher in the greening lakes compared to the blueing lakes (95% CI: 0.11-0.57°C; Figure
197 S4). For lakes that shifted from Blue->Greener, nearly every lake experienced substantial summer

198 warming (Table S2). Precipitation shifts were highly variable, and most lakes did not experience
199 substantial shifts in PRISM-estimated monthly precipitation (Table S2).

200 **Discussion**

201 Our analysis showed that most lakes (55%) included in this study showed no substantial
202 change in lake color between 1984 and 2020. This is consistent with both remote sensing and field
203 studies of regional lake water quality trends in arctic (Kuhn and Butman, 2021) and temperate regions
204 (Oliver et al., 2017; Paltsev and Creed, 2021) that showed a minority of study lakes to be exhibiting
205 changes in lake color. For lakes in the Rocky Mountain region that changed over the past 36 years,
206 most trended bluer (70%), suggesting an overall improvement in summer water quality. While there is
207 a growing concern of widespread declines in water quality, our results build on recent studies that
208 show regional improvements in water quality and a more nuanced understanding of changes in lakes
209 occurring across large spatial scales (Topp et al., 2021; Wilkinson et al., 2021).

210 *Spatial patterns*

211 Our study revealed several putative controls on spatial patterns in lake color in the U.S. Rocky
212 Mountains. Many blue lakes were in steep, high-elevation watersheds, with little vegetative cover and
213 had colder mean annual air temperature than green/brown lakes (Figure S5). Together, these factors
214 likely result in limited terrestrial nutrient subsidies and thus lower productivity and clearer waters
215 (Leavitt et al., 2009; Likens and Bormann, 1974). Heterogeneity in additional factors among these high
216 elevation lakes such as lake morphometry and watershed area may also modify this relationship. For
217 example, some green/brown lakes occurred in cold areas ($MAAT < 4.5^{\circ}C$) if they were shallow ($< 2.5m$
218 average depth), particularly if they had larger watersheds ($> 12.5km^2$). This is expected since small,
219 shallow lakes tend to be more productive than deep lakes (Duarte and Kalff, 1989; Genkai-Kato and
220 Carpenter, 2005; Richardson et al., 2022). Conversely, in some shallow lakes, the color that satellites
221 detect may be capturing benthic algal growth, which can make up a majority of the lake productivity in
222 systems where photic zone extends to the benthos (Lõugas et al., 2020). Overall, these spatial

223 patterns are consistent with studies describing continental scale patterns of lake trophic status and
224 water quality, which indicate that high-elevation western mountain ecoregions are generally
225 oligotrophic, with higher prevalence of green, turbid, or eutrophic lakes in plains and agricultural
226 ecoregions (Hill et al., 2018; Hollister et al., 2016; Peck et al., 2020).

227 *Controls on cross-lake color trends*

228 Lakes and reservoirs that appeared to be improving in water quality (*i.e.* shifted toward bluer
229 wavelengths) represented 32% of all sites and frequently occurred in developed, relatively lower
230 elevation areas. Reservoir management in the Western U.S. typically employs a variety of approaches
231 (e.g., hypolimnetic oxygenation, diversifying water supplies) to maintain water resources under
232 increasing climate variability (Beutel and Horne, 1999; Page and Dilling, 2020; Ray, 2003) and these
233 practices may be a driver of the water quality improvements we observed. However, these apparent
234 changes in water quality that may be attributed to local management actions were difficult to capture
235 in our statistical analyses because we lacked broad-scale databases that summarize management
236 efforts for this region. In addition, managed movement of water across the landscape could further
237 obscure relationships between watershed characteristics and local water quality trends. For example,
238 we observed clusters of reservoirs with blueing trends in the heavily populated Colorado Front Range,
239 but trans-basin water diversions are common in that area (Wiener et al., 2008) making it even more
240 difficult to link management practices to changing water color. Our results suggest that management
241 practices over the same period may have led to improving water quality in ecosystems that are often
242 used for drinking water.

243 A relatively small proportion of lakes (13%) exhibited characteristics indicative of decreasing
244 water quality, either shifting from states of blue to greener or intensifying green/brown. Similarly, recent
245 studies of chlorophyll-a trends in U.S. lakes have shown algal intensification to be occurring in a
246 relatively small proportion of lakes with long-term field data (Wilkinson et al., 2021). Lakes that did
247 exhibit trends toward greener waters were diverse in their size, shape, watershed area, land cover,
248 and climatic changes. This level of spatial heterogeneity has also been shown in regard to

249 cyanobacteria bloom frequency, where the Rocky Mountain region represented a region where blooms
250 were isolated rather than spatially clustered (Coffer et al., 2021). This result reinforces that interactions
251 between local landscape factors and broader climatic changes can result in heterogeneous, context-
252 dependent responses on freshwater systems (Birk et al., 2020; Jackson et al., 2016).

253 Notably, the random forest model had a very limited capacity to classify lakes trending toward
254 greener wavelengths (Figure S2). These greening lakes tended to be at some of the highest elevations
255 and were sparsely populated by humans relative to the lakes that were blueing (Figure 5). Many of
256 these sites experienced slight increases in winter precipitation and decreases in spring temperature.
257 The six lakes that showed the most substantial changes in lake color (Table S1) had very little in
258 common except that they all have experienced increases in mean summer air temperature (1.0-1.95°C
259 since 1984) and were all shallow (less than 3m mean depth), suggesting that lake color in these
260 systems includes bottom reflectance and possible benthic blooms (Vadeboncoeur et al., 2021). While
261 the slope of the greening trends in color in these 39 lakes were statistically significant, we emphasize
262 that most of the color values were within the range of wavelengths that classify these lakes as “blue”
263 following the approach we used in our spatial analysis. Nonetheless, these lakes appear to be on a
264 “greening” trajectory and the underlying cause of that shift warrants further investigation.

265 Winter precipitation and spring temperatures, partially explained temporal trends in summer
266 lake color, but they do not fully capture variability in snowpack regimes (Trujillo and Molotch, 2014).
267 In many mountainous areas, winter and spring snowpacks control the length of ice duration (Caldwell
268 et al., 2021), thus changes in these climatic variables can have cascading effects on lake chemistry
269 and ecology (e.g., algal phenology), and thus color (Cavaliere et al., 2021; Hébert et al., 2021). Less
270 snow in combination with warmer summers may interactively stimulate lake production in systems
271 (Oleksy et al., 2020; Preston et al., 2016), but these same climatic changes can have the opposite
272 effect on lakes in other regions (i.e., lower phytoplankton biomass; Hrycik et al., 2021), highlighting
273 the need to understand how multiple stressors can have either synergistic or antagonistic effects
274 across lakes.

275

276 Finally, there are a few possible explanations for why we did not detect widespread changes
277 in lake color in the region. First, our dataset only included relatively large lakes that were ≥ 10 ha, but
278 most lakes in the Rocky Mountains are < 10 ha (85.3%, $n=15,568$) and the smallest lakes are more
279 abundant at high elevations (Figure S5). This may partially explain why rates of nitrogen deposition
280 did not appear to have an effect of water color trends, even though excess nitrogen is implicated as a
281 driver of ecological change in high-elevation lakes across the region (Burpee et al., 2022; Moser et
282 al., 2019; Oleksy et al., 2020). Second, we limited our analysis to median summer color, but it is
283 possible that there are dynamics that have helped create the perception of lake greening, such as
284 episodic algal blooms, which are increasing in some systems (Ho et al., 2019; Vadeboncoeur et al.,
285 2021; Wilkinson et al., 2021). This could create issues where algal blooms really are present but are
286 short and intense and thus not captured by Landsat's 8 or 16 day return sampling interval. As such,
287 algal blooms that are increasing in severity, duration, or magnitude may not be detected by our
288 approach. Conversely, by limiting our analysis to summer months, we may be missing shifts in the
289 phenology of lake color, such as early greening in the spring or a second peak of productivity in the
290 fall (Sommer et al., 2012). Future studies related to lake changes may consider changes and variability
291 in the entirety of the ice-free season. Furthermore, our understanding of regional changes in water
292 quality will be greatly enhanced by advances in the remote sensing of small lakes.

293 **Conclusions**

294 Climate change impacts are likely to influence high-elevation systems faster than others, making high-
295 elevation lakes sentinels of climate change (Adrian et al., 2009; Moser et al., 2019). While
296 eutrophication could pose a major threat to the ability for these systems to continue to provide their
297 vital services to downstream communities, we found that the majority of large lakes (> 10 ha) in this
298 region are not changing. Where we did observe lake color changing, it was consistently towards bluer
299 waters. However, some of the mechanisms for the observed changes, particularly in greening lakes,
300 remain elusive. Future work in this region should investigate how the slow, press changes from climate

301 change interact with short, intense pulse disturbances like floods and fire to alter the ecology of Rocky
302 Mountain lakes and reservoirs.

303 **Conflict of Interest Statement**

304 The authors declare no conflicts of interest relevant to this study.

305 **Acknowledgement**

306 This work was supported by NSF award #EPS-2019528.

307 **Data Availability Statement**

308 The data for this paper comes from the Landsat Archive (via LimnoSat-US
309 10.5281/zenodo.4139695), LAGOS-US Reservoirs, LAGOS-US NETWORKS, EPA's LakeCat,
310 and The Global Lake Area, Climate, and Population database. All these data are free to
311 download with appropriate links in the bibliography.

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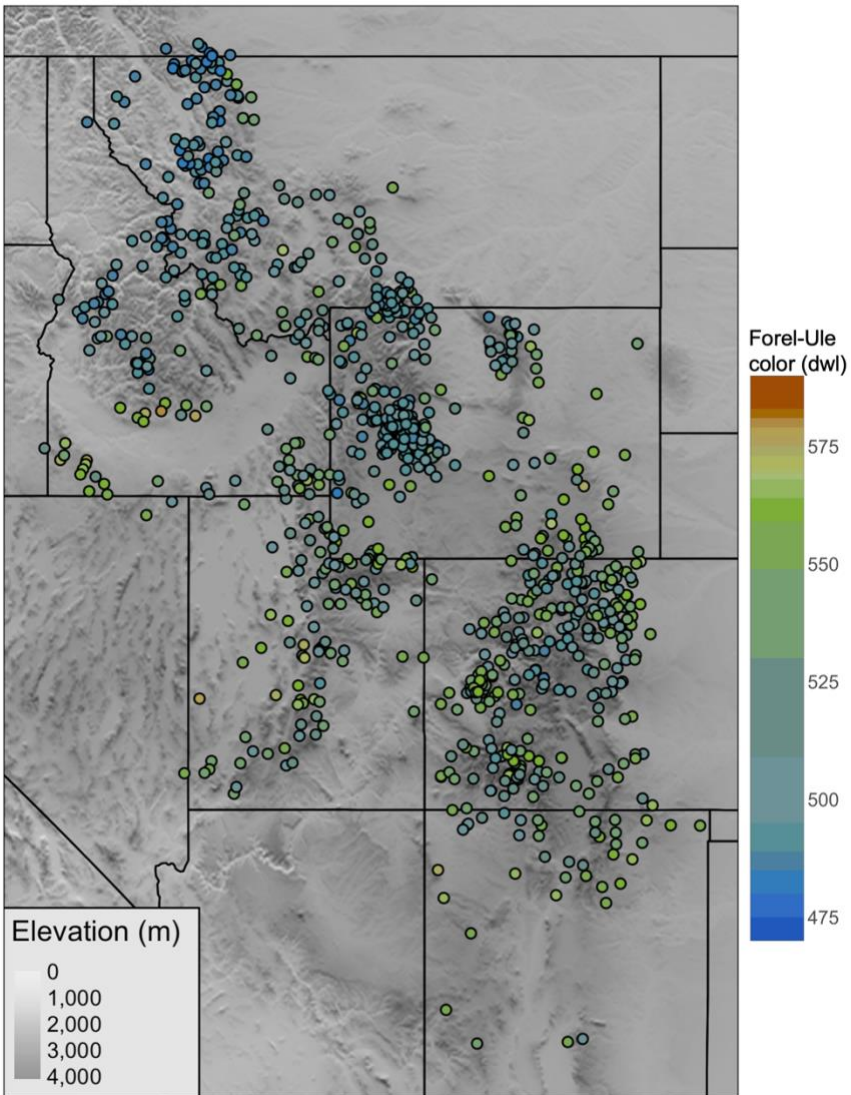
524

525 **Table 1.** Covariates included in the spatial CART model (*) and the temporal Random Forest model (†).

Variable	Mean (sd)	Description	Data source
precip.	47.8 (14.2)	mean monthly precipitation (mm)*	Labou et al., 2020
air temp.	4 (2.8)	mean annual air temperature (°C)*	Labou et al., 2020
population	14,182.4 (166,828.8)	total human population*†	GLCP
LA (km ²)	2.6 (12.9)	lake surface area (km ²)*†	NHD
WA:LA	231.8 (894.6)	watershed area:lake area	
WSA	475.7 (2537)	watershed area (km ²) *†	NHD
elev.	2,291.9 (571.9)	lake elevation (m) *†	NHD
Z _{mean}	4.8 (4.8)	mean lake depth (m) *†	NHD
Z _{max}	12.5 (13.3)	max lake depth (m) *†	NHD
NO ₃ dep.	3.3 (1.4)	total nitrate deposition (2018) *†	NADP
NH ₃ dep.	1.9 (0.9)	total ammonia deposition (2018) *†	NADP
% ice	0.3 (1.4)	% watershed area classified as ice/snow land cover*†	NLCD
% urban	0.7 (4.8)	% watershed area classified as developed, low+med+high-intensity land use*†	NLCD
% forest	3.7 (9.8)	% watershed area classified as deciduous, coniferous, and mixed forest land cover*†	NLCD
% shrub	31.3 (23.7)	% watershed area classified as shrub/scrub land cover	NLCD
% grassland	18.6 (21.3)	% watershed area classified as grassland/herbaceous land cover	NLCD
% agriculture	2.1 (8.7)	% watershed area classified as crop and hay lake cover	NLCD
% wetland	1.9 (4.7)	% watershed area classified as herbaceous+woody wetland land cover	NLCD
% barren	4.1 (8.6)	% watershed area classified as barren land cover	NLCD

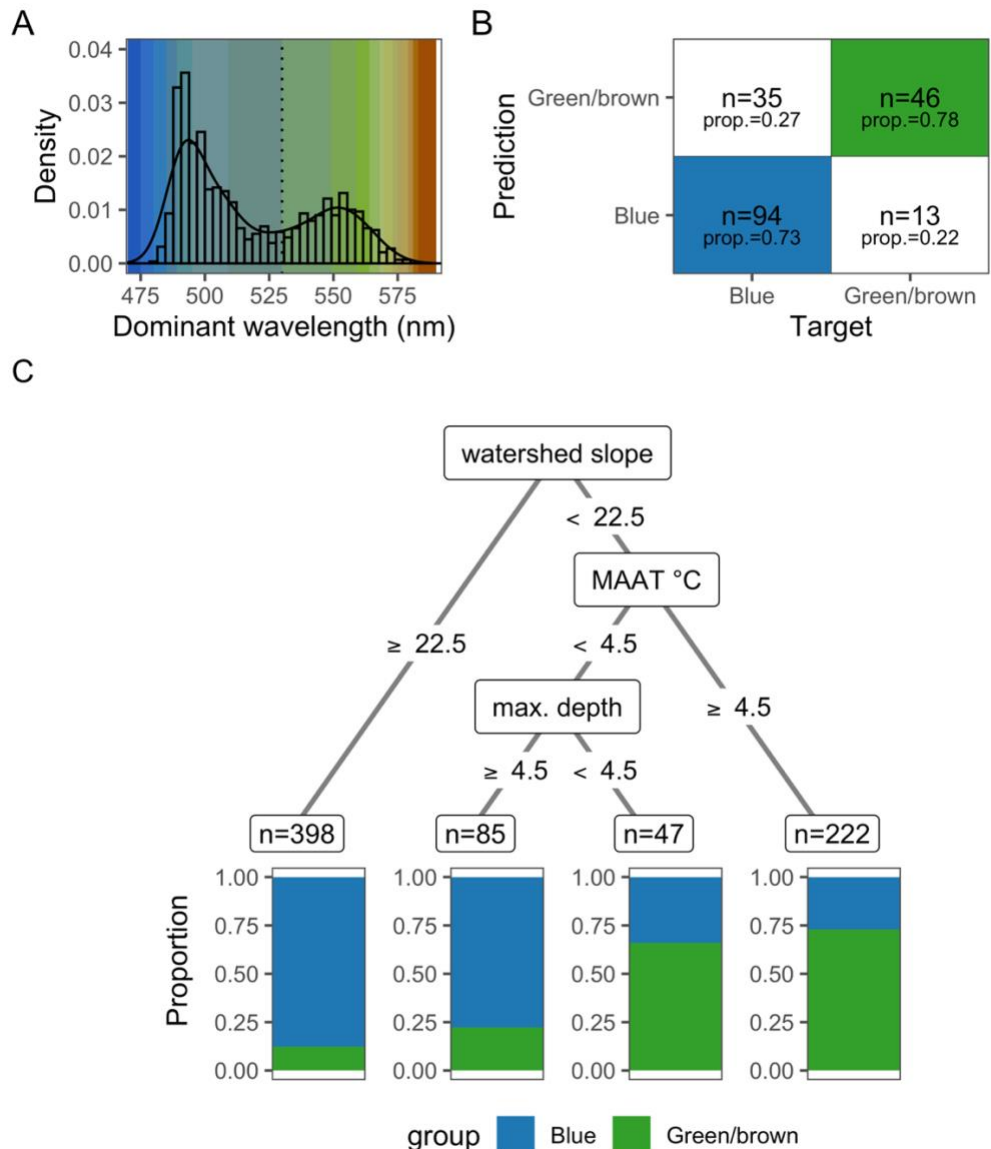
carb.	4 (14.6)	carbonate bedrock*†	LakeCat
sil.	46.9 (44.2)	silicate bedrock*†	LakeCat
slope	25.8 (16)	mean watershed slope angle	LakeCat
CTI	734.7 (110)	mean Composite Topographic Index (CTI) within catchment	LakeCat

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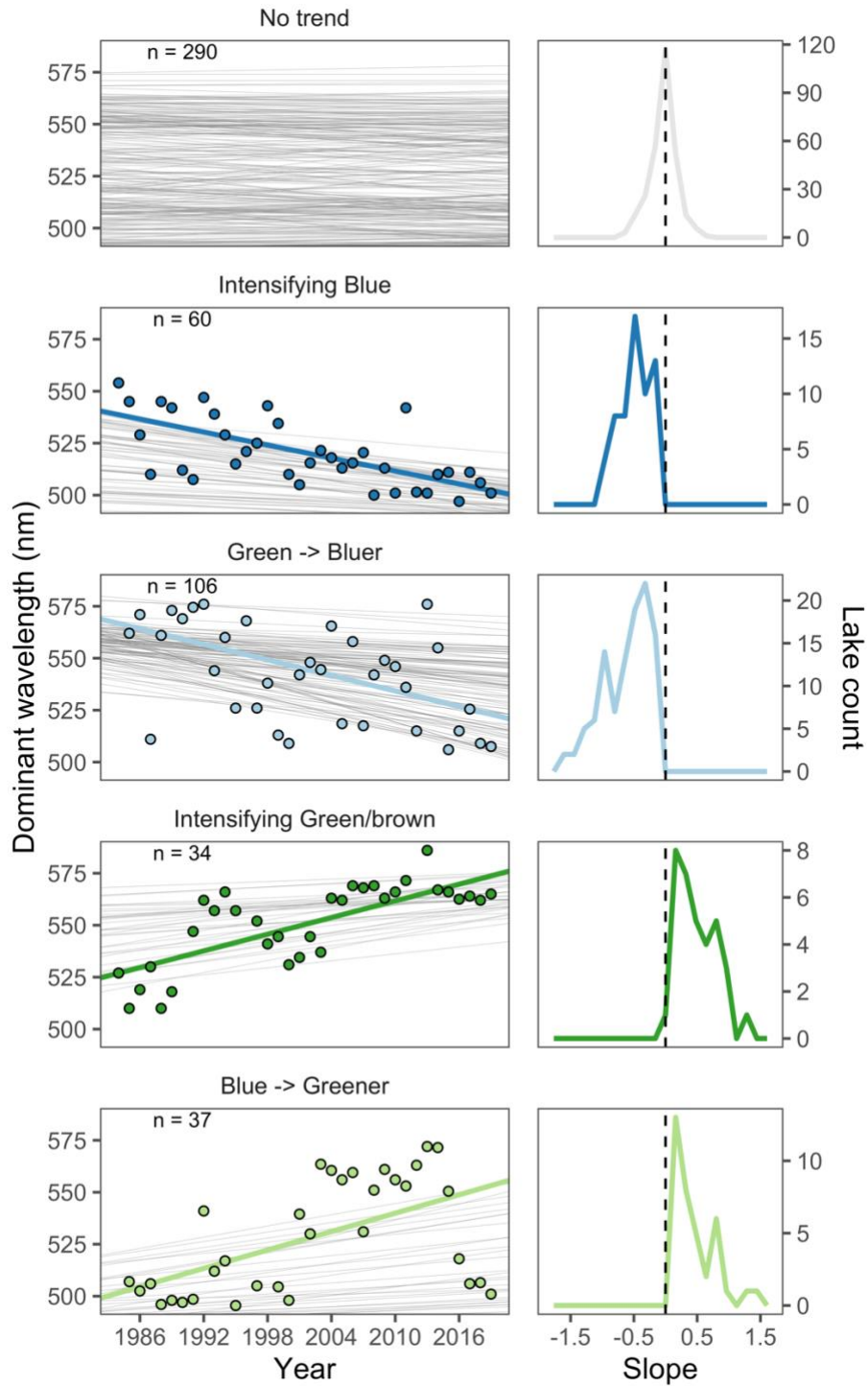
527

528 **Figure 1:** Spatial distribution of median lake color for the study lakes from the modern (2010-2020)
 529 period. Lakes are shaded by the 2010-2020 median dominant wavelength (DWL) and its
 530 corresponding color on the Forel-Ule scale. Note that individual points are jittered so that points with
 531 similar values do not overlap, therefore lake locations are approximate.



532

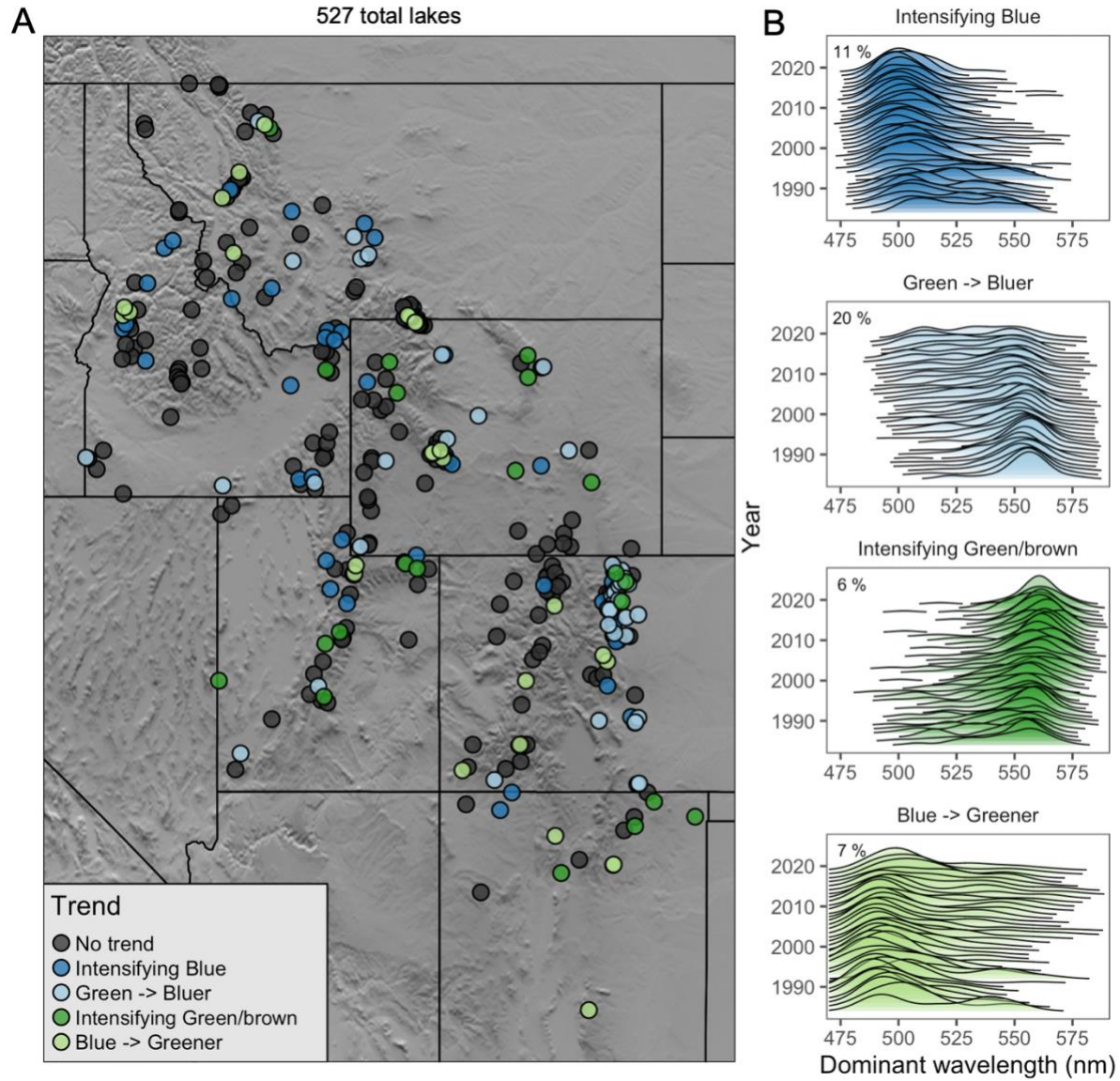
533 **Figure 2: A.** Density plot showing the distribution of a median dominant wavelength (2010-2020)
 534 where the background color corresponds directly to the Forel-Ule index color. Vertical dashed line
 535 represents our threshold for classifying lakes as blue vs. green. **B.** Confusion matrix for the testing
 536 data of the spatial CART where true positives for blue/clear classifications are shaded in blue and true
 537 positives for green/murky lakes are shaded in green; depicts the accuracy when assigning blue or
 538 green lake grouping to a set of lakes that were not used in the training algorithm **C.** CART model
 539 results are visualized in tree form, where the terminal node shows the proportion of blue or
 540 green/brown lakes.



541

542 **Figure 3.** Example trends on the left for each trend category with the corresponding frequency

543 polygons of calculated Sen's slopes on the right.



544

545 **Figure 4. A.** Regional map on the left shows where lakes fall into trend categories. **B.** Panels show

546 distributions of dominant wavelength through time in each of the changing trend categories.

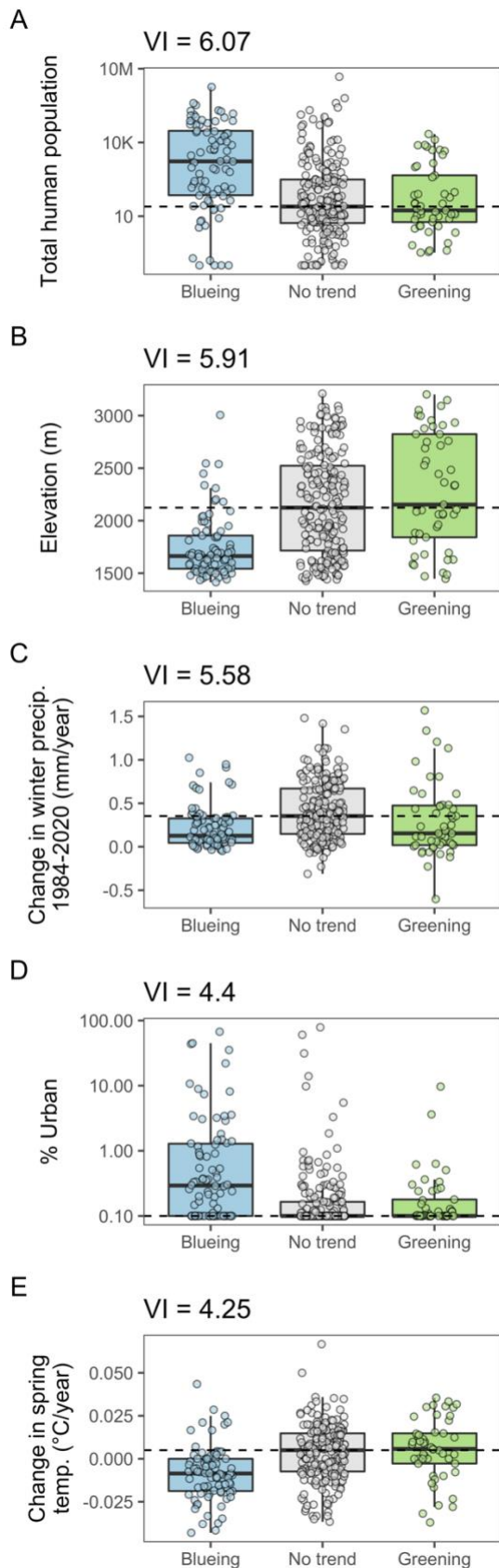


Figure 5 (left). The top 5 variables from highest to lowest based on mean decrease in accuracy from the Random Forest temporal trend classification model. The dashed line shown on each figure is to aid the reader in comparing the different trend categories to the median value for the “No Trend” lakes. Lakes that changed from Green -> Blue or were Intensifying Blue had negative trends in dominant wavelength while lakes that changes from blue -> greener or were Intensifying Green/brown had positive trends in dominant wavelength.

1

Supplementary Information for

2 **Title:** Heterogenous controls on lake color and trends across the high-elevation U.S. Rocky Mountain
3 region

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5 Hall^c, Catherine M. O'Reilly^d, Xiao Yang^b, Matthew R.V. Ross^c

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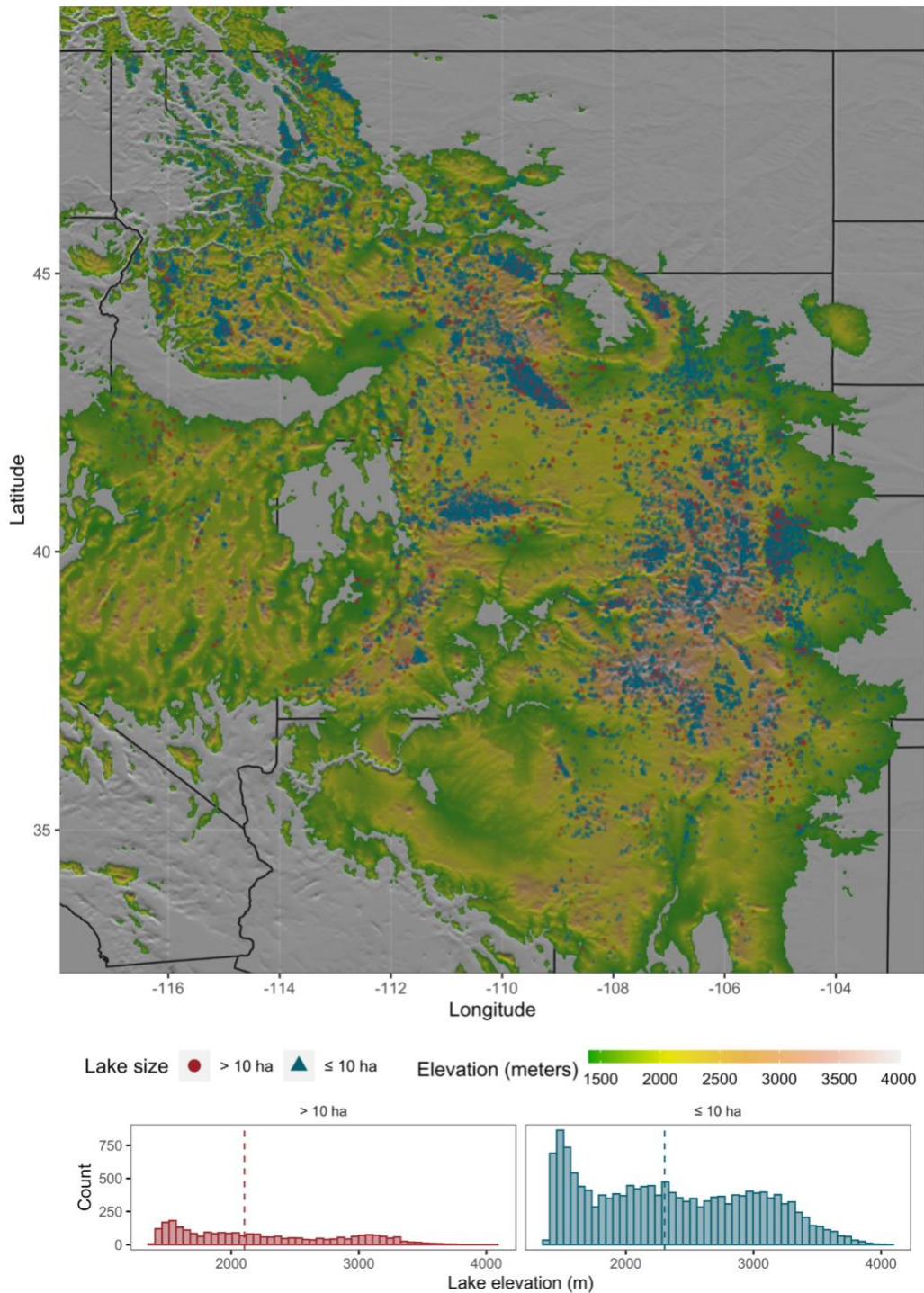
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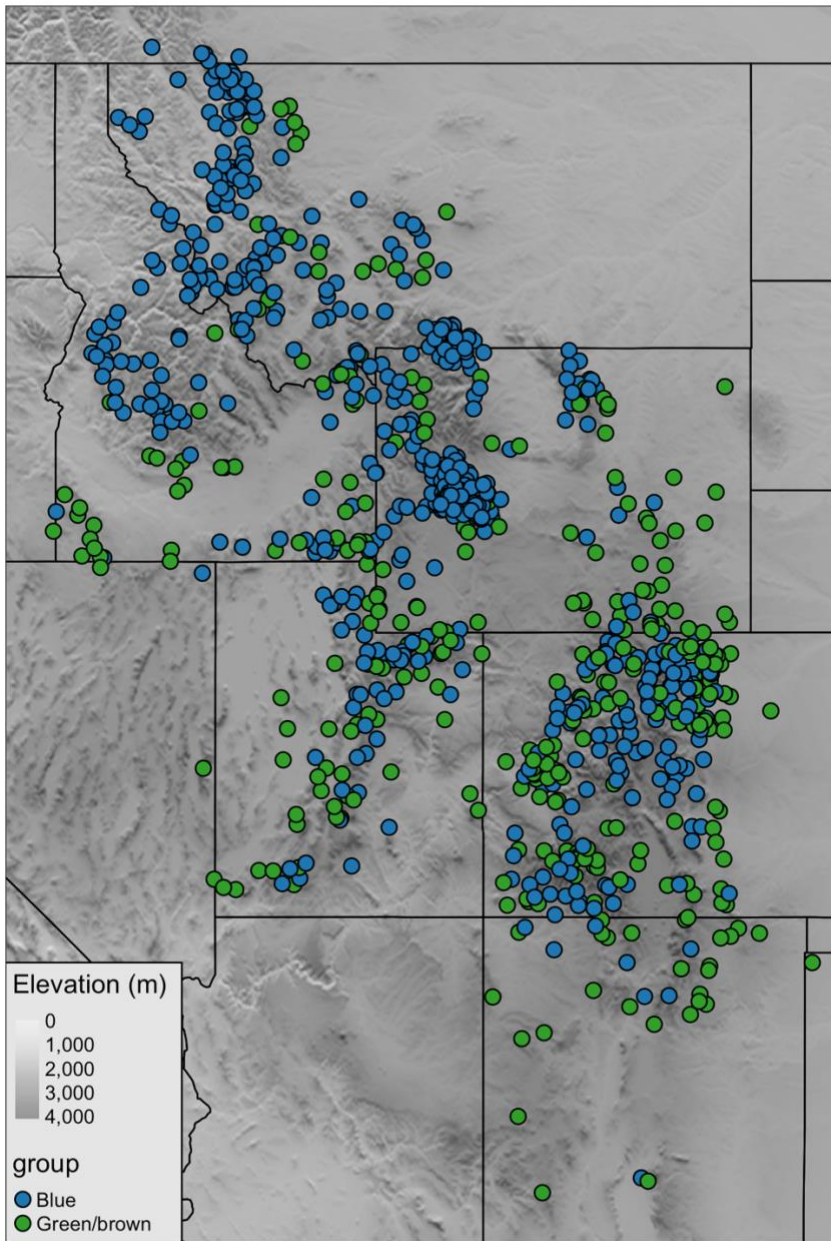
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13 USA

All lakes >1 ha and >1400m in the study region
 There are 18,256 lakes total, and only 14.7% (n=2688) are > 10 ha in size



14

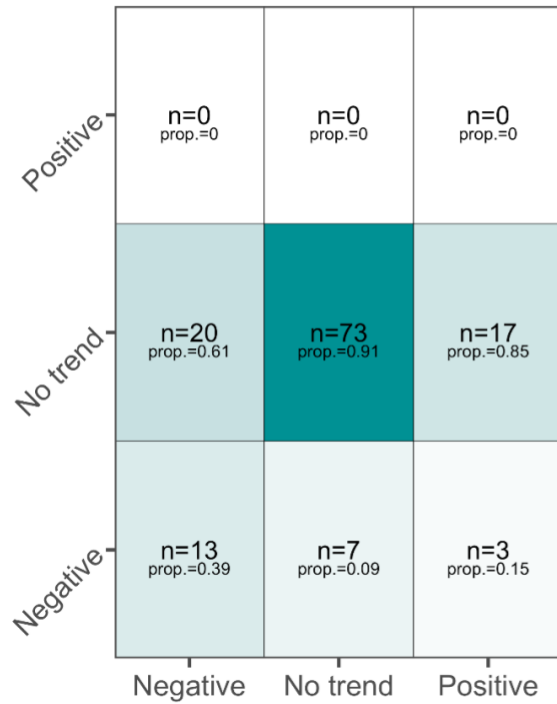
15 **Figure S1.** Map of all lakes greater than 1 ha and over 1400 meters elevation in the greater
 16 Rocky Mountain region of the United States. Dashed lines on the histograms show the median
 17 elevation for each lake size class. Lake size and location from (Smith et al., 2021).



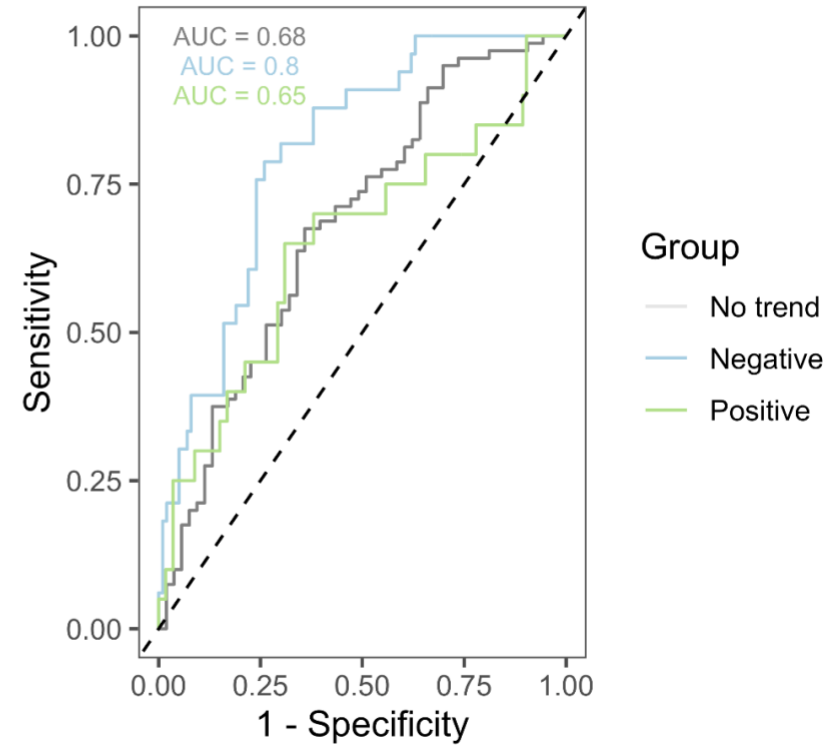
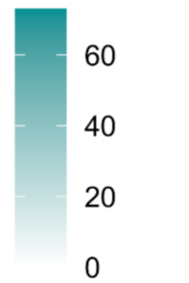
19

20 **Figure S2.** Spatial distribution of median lake color for the study lakes from the modern period
 21 (2010-2020) based on a binary classification scheme. Lakes with median dominant wavelength
 22 greater than 530 nm were classified as Green/brown while lakes less than 530 nm were
 23 classified as generally Blue.

Accuracy: 0.65

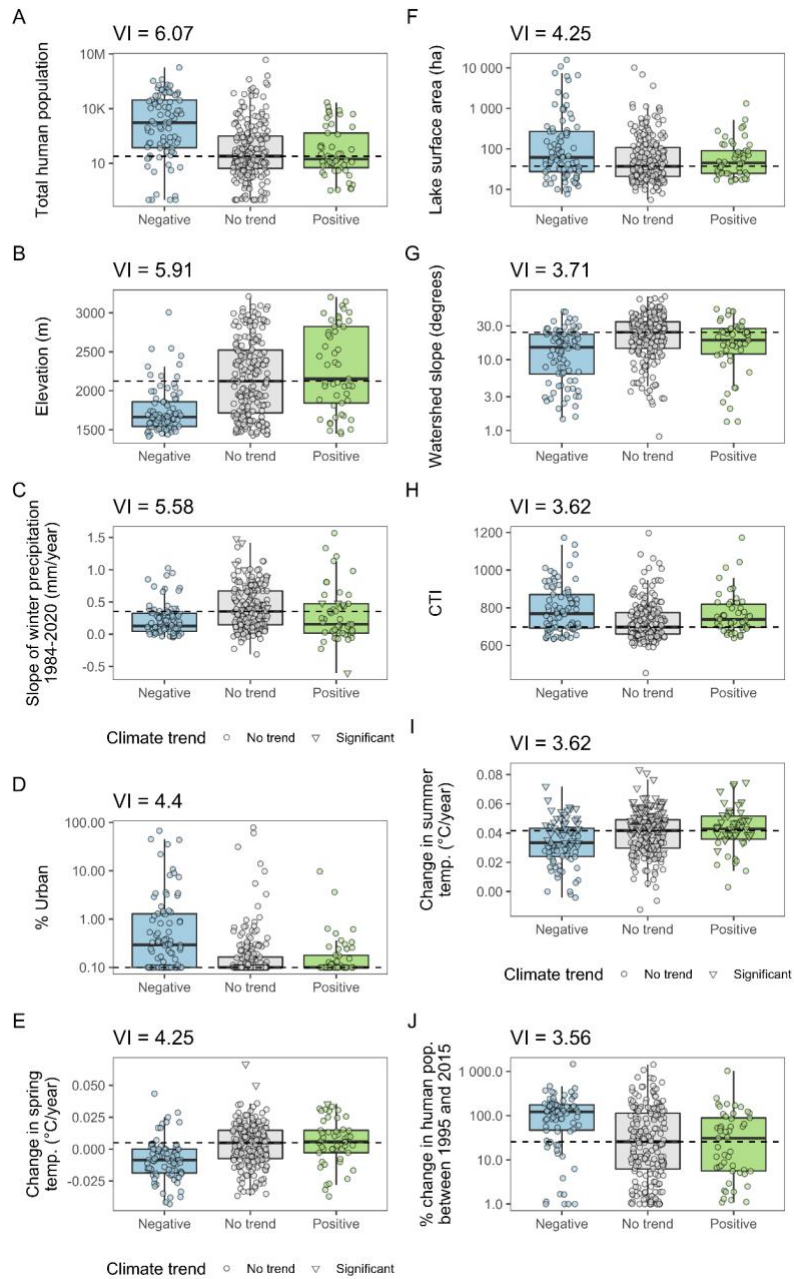


Frequency



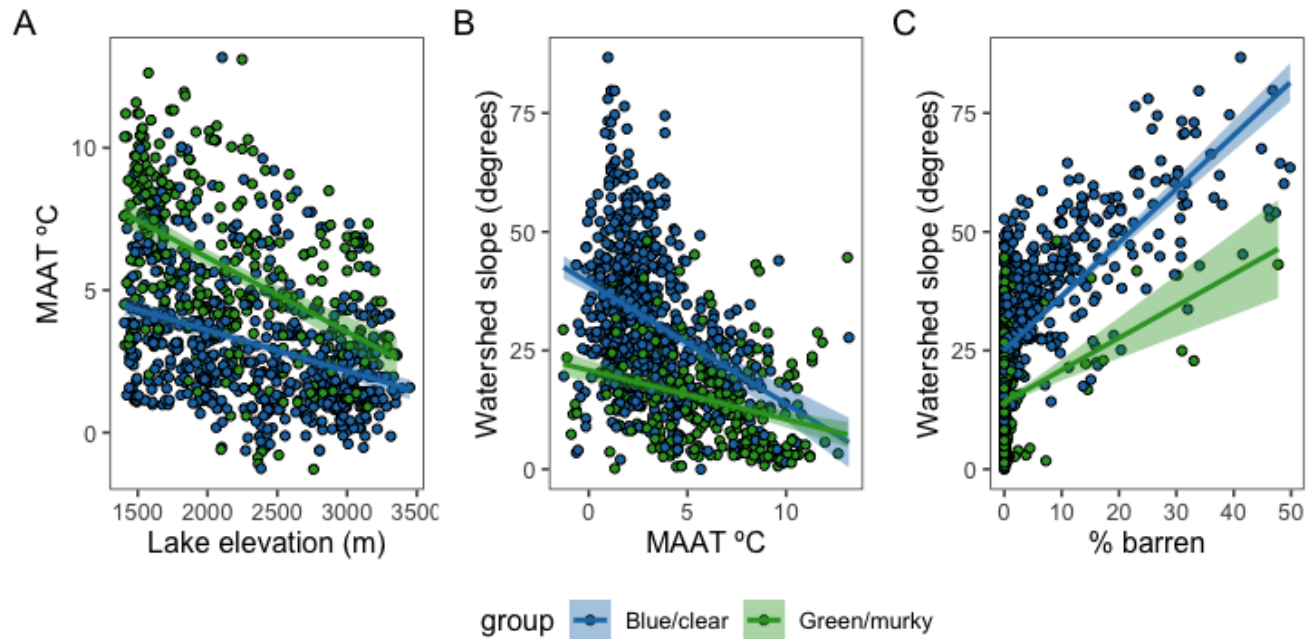
24

25 **Figure S3.** Confusion matrix and ROC curve for the random forest trend classification model.



26

27 **Figure S4.** The top 10 variables from highest to lowest based on mean decrease in accuracy
 28 from the Random Forest temporal trend classification model. The dashed line shown on each
 29 figure is to aid the reader in comparing the different trend categories to the median value for the
 30 “No Trend” lakes. In panels C, E, and I the shape of the points denotes whether the Sen’s
 31 slopes on the winter precipitation, spring, and summer temperatures trends are statistically
 32 significant.



33

34 **Figure S5.** Relationships among key predictors of spatial variability in lake color. (A) Mean annual temperature is negatively
 35 correlated with lake elevation. (B) Watershed slope is negatively correlated with mean annual air temperature. (C) Watershed slope
 36 is positively correlated with % barren landcover.

37 **Table S1.** The subset of lakes in the Blue -> Greener and Green -> Bluer categories which last color shifts such that they were
 38 classified as either Green/brown or Blue in the spatial analysis even though early in the record (1984-2005) these lakes would have
 39 been classified as Blue or Green/brown (respectively).

Trend	Mode change	Hylak ID	Lake name	State	Lake type	Lat	Lon
Blue -> Greener	Blue to Green/brown	995352	Mitten Lake	MT	Natural lake	-113	48.3
Blue -> Greener	Blue to Green/brown	1060135	Lake David	NM	Natural lake	-105	35.8
Blue -> Greener	Blue to Green/brown	112872	Lake Isabel	NM	Natural lake	-105	35.8
Blue -> Greener	Blue to Green/brown	1055853	Hinman Reservoir	CO	Reservoir	-106	40.2
Blue -> Greener	Blue to Green/brown	1058576	Hermit Lakes	CO	Reservoir	-107	37.8
Blue -> Greener	Blue to Green/brown	1059035	Totten Lake	CO	Reservoir	-109	37.4
Green -> Bluer	Green/brown to Blue	1056066	Unnamed	CO	Reservoir	-105	40
Green -> Bluer	Green/brown to Blue	1055647	Ryan Gulch Reservoir	CO	Reservoir	-105	40.4
Green -> Bluer	Green/brown to Blue	1055851	Clover Basin Reservoir	CO	Reservoir	-105	40.2
Green -> Bluer	Green/brown to Blue	1055926	Left Hand Valley Reservoir	CO	Reservoir	-105	40.1
Green -> Bluer	Green/brown to Blue	1056206	Ralston Reservoir	CO	Reservoir	-105	39.8
Green -> Bluer	Green/brown to Blue	112156	Boulder Reservoir	CO	Reservoir	-105	40.1
Green -> Bluer	Green/brown to Blue	112721	Trinidad Lake	CO	Reservoir	-105	37.1
Green -> Bluer	Green/brown to Blue	1055703	Hertha Reservoir	CO	Reservoir	-105	40.3
Green -> Bluer	Green/brown to Blue	111430	Unnamed	ID	Reservoir	-114	42.2
Green -> Bluer	Green/brown to Blue	1050316	Dougal Reservoir	ID	Reservoir	-117	42.7
Green -> Bluer	Green/brown to Blue	1027810	Wertz Reservoir	MT	Reservoir	-111	46.4
Green -> Bluer	Green/brown to Blue	1031268	Unnamed	MT	Reservoir	-111	46.1
Green -> Bluer	Green/brown to Blue	1050483	McNinch Number 1 Reservoir	WY	Reservoir	-110	42.6
Green -> Bluer	Green/brown to Blue	9066	Boysen Reservoir	WY	Reservoir	-108	43.4

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41 **Table S2.** For each lake-year, we calculated the mean winter (December-February), spring (March-May), summer (June-August), and fall
 42 (September-November) temperature and precipitation trends from PRISM (see *Methods* for details). We calculated the non-parametric
 43 Theil-Sen's slope for each lake time series of season precipitation or temperature. In the tables below, we summarized the number of lakes
 44 (and percentage of the lakes in each Trend-Model shift combination) that showed substantial trends in precipitation or temperature, using
 45 the Mann-Kendall z-score to test for statistical significance at the $\alpha = 0.05$ level. Dashes (-) indicate that no lakes showed trends in that
 46 particular category. Overall, the most widespread climatic trends in the region were increasing summer and fall temperatures. Lakes with
 47 color trends classified as Blue -> Greener, nearly every lake has experienced substantially summer warming. Precipitation trends were
 48 much more variable, and most lakes have not experienced large shifts in PRISM-estimated monthly precipitation.
 49

Sens slope Trend direction			Winter temps.		Spring temps.		Summer temps.		Fall temps.	
			↑	↓	↑	↓	↑	↓	↑	↓
No trend	No trend	No net change	-	-	-	1 (0.3%)	182 (62.8%)	-	206 (71%)	-
Negative	Intensifying Blue	No net change	-	-	-	-	39 (65%)	-	35 (58.3%)	-
Negative	Green -> Bluer	Green/brown to Blue	-	-	-	-	11 (73.3%)	-	10 (66.7%)	-
Negative	Green -> Bluer	No net change	-	-	-	-	56 (62.2%)	-	78 (86.7%)	-
Positive	Intensifying Green/brown	No net change	-	-	-	-	17 (51.5%)	-	28 (84.8%)	-
Positive	Blue -> Greener	No net change	-	-	-	-	27 (87.1%)	-	25 (80.6%)	-
Positive	Blue -> Greener	Blue to Green/brown	-	-	-	-	6 (100%)	-	5 (83.3%)	-

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Sens slope Trend direction			Winter precip.		Spring precip.		Summer precip.		Fall precip.	
			↑	↓	↑	↓	↑	↓	↑	↓
No trend	No trend	No net change	33 (11.4%)	-	11 (3.8%)	5 (1.7%)	-	16 (5.5%)	2 (0.7%)	-
Negative	Intensifying Blue	No net change	7 (11.7%)	-	3 (5%)	2 (3.3%)	-	-	-	-
Negative	Green -> Bluer	Green/brown to Blue	1 (6.7%)	-	1 (6.7%)	-	-	-	-	-
Negative	Green -> Bluer	No net change	6 (6.7%)	-	9 (10%)	-	-	2 (2.2%)	-	-
Positive	Intensifying Green/brown	No net change	6 (18.2%)	-	3 (9.1%)	-	-	3 (9.1%)	-	-
Positive	Blue -> Greener	Blue to Green/brown	3 (50%)	-	-	-	-	-	-	-
Positive	Blue -> Greener	No net change	1 (3.2%)	1 (3.2%)	-	-	-	2 (6.5%)	-	-

51

52 **References**

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