

Mechanistic insights into sulfur rich oil formation, relevant to geological carbon storage routes. A study using (+) APPI FTICR-MS analysis

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1 **Mechanistic insights into sulfur rich oil formation, relevant to geological carbon storage**
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7
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9 sequestration

10

11 **Abstract**

12 Sulfur incorporation into sedimentary organic matter has a key role in carbon preservation in the
13 geosphere. Such processes can inform strategies for human timescale carbon storage to mitigate
14 climate change impacts and thus more detailed knowledge of sulfur incorporation into biomass
15 species is needed. Until recently, detailed chemical characterization of sulfurized organic matter
16 was only possible by analyzing individual building blocks obtained after desulfurization
17 reactions. In this study, Fourier transform ion cyclotron resonance mass spectrometry (FTICR-
18 MS), with atmospheric pressure photoionization in positive ion mode, (+) APPI, was used to
19 investigate the chemical composition of sulfur rich crude oils and to obtain mechanistic insights
20 into the sulfur incorporation reactions happening during early diagenesis. Contrary to
21 expectations, (+) APPI FTICR-MS data show that sulfurized lipids (with up to 6 sulfur atoms
22 and up to m/z 1100) occur as free molecules in these oils, rather than within a macromolecular
23 network linked by (poly)sulfide bridges. In contrast to the mature Peace River (Canada) oils, the
24 thermally immature Rozel Point (USA) and Jiangnan Basin (China) oils show a carbon number
25 preference in sulfurized species resembling biogenic precursor molecules, which highlights the
26 importance of S-bound molecules as geochemical proxies for early diagenetic processes. This
27 study indicates that sulfur incorporation reactions involve the formation of S-cyclic structures in
28 which the double bond equivalent is \geq the number of S atoms. Collision induced dissociation

29 (CID-) FTICR-MS experiments suggest the occurrence of intermolecular sulfur incorporation
30 reactions, but only as a mechanism that is secondary to intramolecular sulfur addition. The CID-
31 FTICR-MS experiments indicated that steroid sulfurization typically yields S-bearing cyclic
32 structures and that thiol/thioether groups may be present throughout the chemical matrix but only
33 to a minor extent. In addition, CID-FTICR-MS also confirms the occurrence of sulfurized
34 alkenones in low maturity oils. Knowledge of organic sulfur molecule formation informs routes
35 for carbon dioxide removal technologies that could be used to sequester carbon in the geosphere
36 and/or hydrosphere in the form of recalcitrant organic species.

37

38 **1. Introduction**

39 Sulfur incorporation into sedimentary organic matter has a key role in carbon preservation in the
40 geosphere, not only on Earth but also possibly on Mars (Werne et al., 2004; Summons et al.,
41 2011). The understanding of mechanisms and geochemical implications of sulfur incorporation
42 reactions expanded greatly during the late 1980s and early 1990s (Orr and Sinninghe Damsté,
43 1990; Aizenshtat et al., 1995; Amrani, 2014), and given their relevant role in enabling an
44 important carbon sink, there are still many active areas of research in the field. For example,
45 there currently exists a heightened research interest to further understand carbon preservation
46 mechanisms via sulfur incorporation (Raven et al., 2016; Pohlabein et al., 2017) in the context of
47 the global change and the need to keep the global temperature increase to <2 °C by the end of
48 this century (IPCC, 2018).

49 Large scales of carbon removal are not unusual in natural systems. The global ocean has pools of
50 organic carbon found in dissolved and particulate form in the water column, or deposited in the
51 deep-sea sediments that are of similar magnitude to atmospheric carbon reservoirs (Houghton,
52 2007; Hansell et al., 2009; Jiao et al., 2010). Therefore, a better understanding of natural
53 mechanisms for long term preservation of organic molecules in the geosphere can provide
54 guidelines for the development of technologies that artificially enhance the removal of carbon
55 from the atmosphere to ocean or subsurface reservoirs, in the form of biologically refractory
56 organic species analogous to those already present naturally in the Earth. Kinetics and feedback
57 mechanisms of any engineered interventions within natural organic matter pools will be crucial
58 variables to achieve net negative carbon drawdown from the atmosphere on the timescales
59 relevant for anthropogenic climate intervention (hundreds to thousands of years).

60 During the early diagenesis, sulfurized organic compounds, commonly occurring in anoxic
61 environments that are rich in reduced sulfur, typically resemble their biological precursors. As
62 the diagenesis progresses, extensive sulfurization reactions enable multi-point, intermolecular
63 cross-linkage, eventually generating the macromolecular structures that are the building blocks
64 of sulfur rich kerogen (type II-S) (Tegelaar et al., 1989; Vandenbroucke and Largeau, 2007).
65 Sulfur rich macromolecular structures were also detected in thermally immature, sulfur rich oils
66 (Sinninghe Damsté et al., 1987; Sinninghe Damsté et al., 1989; Orr and Sinninghe Damsté,

67 1990). However, the detailed chemical characterization of such compounds was only possible by
68 analyzing individual building blocks after desulfurization reactions (Adam et al., 1992; Adam et
69 al., 1993; Richnow et al., 1993). Since then, analytical technologies have greatly advanced.
70 Ultrahigh resolution mass spectrometry, in particular the FTICR-MS (Fourier transform ion
71 cyclotron resonance mass spectrometry) technology, has enabled the identification of several
72 thousand peaks in complex organic mixtures, revolutionizing the understanding of the
73 petroleome (Marshall and Rodgers, 2008). In the field of organic geochemistry, such technical
74 advances were translated into an unprecedented capability to probe the high molecular weight,
75 polar compounds at a high level of detail, expanding the way biodegradation, thermal stress and
76 other biogeochemical processes are understood (Oldenburg et al., 2014; Radović et al., 2015;
77 Radović et al., 2016b; Oldenburg et al., 2017).

78 Investigations focused on the geochemical significance of sulfur species have also greatly
79 benefited from ultrahigh resolution mass spectrometry applications. Hughey et al. (2004)
80 analyzed two same sourced Smackover Formation oils of different levels of thermal maturity
81 using (-) ESI (electrospray ionization in negative ion mode) FTICR-MS, suggesting the removal
82 of sulfur and oxygen compounds was promoted by thermal maturation. Similarly, Oldenburg et
83 al. (2014) reported that the relative apparent abundances of all heteroatom-containing compound
84 classes (nitrogen, oxygen, sulfur and mixed heteroatom species) detected in that study decreased
85 systematically with increasing oil maturation levels. Walters et al. (2015) used both (+) APPI
86 (atmospheric pressure photoionization in positive-ion mode) and (-) ESI FTICR-MS to probe
87 sulfur compounds and oxygenated analogs present in samples from the Smackover Formation,
88 revealing intermediates (classes O_x and SO_x) of thermochemical sulfate reduction (TSR) redox
89 reactions. A similar approach detected the formation of TSR induced proto-solid bitumen,
90 including highly condensed polynuclear aromatic and naphthenoaromatic species with up to
91 three sulfur atoms (Walters et al., 2011). Lu et al. (2013) used (+) ESI (electrospray ionization in
92 positive-ion mode) FTICR-MS to investigate sulfur rich heavy oils from Bohai Bay Basin,
93 reporting a wide range of sulfurized sedimentary steroids in addition to a complex distribution of
94 sulfur and oxidized sulfur compounds. Based on a similar approach, Lu et al. (2014) detected
95 high levels of alkylcyclothioethers within Jiangnan Basin oils. More recently, Liu et al. (2018)
96 reported an in-depth investigation of polar sulfur compounds in immature crude oils from the

97 Jiangnan Basin via (-) ESI and (+) ESI FTICR-MS showing that the detected S-bearing species
98 are mostly cyclic, originating from intramolecular sulfurization of functionalized precursors
99 during early diagenesis. The authors inferred that the extent of intramolecular sulfurization is
100 based on the number of reactive functional groups in the precursor molecule.

101 Expanding on these findings, the present study investigates the compounds present in a suite of
102 sulfur rich oils and their fractions via (+) APPI FTICR-MS and (+) APPI CID-FTICR-MS
103 experiments. (+) APPI is the ionization technique of choice for sulfur-rich sample analysis due to
104 its enhanced ability to ionize sulfur compounds in complex organic mixtures without chemical
105 derivatization or special sample preparation strategies (Purcell et al., 2006; Oldenburg et al.,
106 2014). Extended compositional information of high molecular weight, sulfur bearing molecules
107 that could not be detected in earlier GC based studies are presented and critically discussed based
108 on the current understanding of sulfur incorporation into organic matter. This paper aims to
109 gather new insights from sulfur rich oil molecular composition to probe the mechanisms of sulfur
110 incorporation into sedimentary organic matter, which could possibly be leveraged for the
111 development of carbon dioxide removal (CDR) technologies.

112 **2. Materials and methods**

113 *2.1. Sample set description*

114 Five sulfur rich oils from different basins were selected in this study: (a) three high maturity,
115 highly biodegraded oils from the Peace River area, Canada; (b) one immature oil from the
116 Jiangnan Basin, Eastern China; and (c) one immature oil from the Rozel Point, Utah, USA
117 (Table 1).

118 Rozel Point is an immature, sulfur-rich heavy oil generated from a hypersaline, lacustrine source
119 rock of Miocene age (Meissner et al., 1984; ten Haven et al., 1988; Sinninghe Damsté et al.,
120 1989). High concentrations of organic sulfur (up to 15 wt%) and the unique depositional
121 environment that characterizes Rozel Point oil have motivated many studies of the origin and
122 fate of organically bonded sulfur during early diagenesis, with implications for the understanding
123 of Type II-S kerogen formation (Eglinton et al., 1994). The Jiangnan Basin, located in eastern
124 China, is comprised of five major tectonic units with most oil production occurring in the
125 Qianjiang depression from the Eocene Qianjiang Formation (Philp and Zhaoan, 1987). Oils from

126 the Qianjiang depression are generated from the Qianjiang Formation and Xingouzui Formation,
127 which were deposited in anoxic, sulfate reducing and saline lacustrine environments during the
128 early Cretaceous to Paleogene (Carroll and Bohacs, 2001; Hou et al., 2017). Due to cycles of
129 marine transgression and progradation activity, hot paleoclimate, and clastic sediment deposition
130 from lacustrine systems, over 220 evaporitic layers interlaced with shales and sandstones formed
131 the Jiangnan Basin source rocks (Philp and Zhaoan, 1987; Peirong et al., 2008; Hou et al., 2017).
132 The Qianjiang Formation acts as both source and reservoir (Philp and Zhaoan, 1987; Huang and
133 Hinnov, 2014; Hou et al., 2017). Upper Qianjiang sections act as the oil reservoir while the
134 deeper sections and Xingouzui Formation embody the source rock. Oils from the Qianjiang
135 Formation typically exhibit a pronounced even/odd *n*-alkane predominance, low ratios of Pr/Ph
136 and high ratios of gammacerane/hopane, indicating a highly anoxic, reducing and saline
137 depositional environment (Philp and Zhaoan, 1987). The Peace River oil sands, located in
138 northwestern Alberta, is one of three major bitumen deposits in that province. The Peace River
139 area is comprised of two main groups, the Bullhead Group formed by the Cadomin and Gething
140 formations, and the overlying Fort St. John Group which, in succession, comprises the Bluesky,
141 Spirit River, Peace River and Shaftesbury formations. Peace River oil reservoirs, Bluesky,
142 Gething and McMurray, are late Paleozoic to Mesozoic in age and are charged by Cretaceous oil
143 source rocks (Adams et al., 2012). The reservoirs are supplied with a mixture of hydrocarbons
144 expelled from multiple source rocks including the Nordegg Member of the Fernie Formation,
145 Gordondale Formation and Permian Doig Formation (Riediger, 1994; Adams et al., 2013).
146 Manville Group Bluesky and Gething reservoirs may also be vertically charged by the
147 Mississippian Exshaw-Banff formations where hydrocarbons migrated upwards into permeable
148 Cretaceous sand units when erosion deteriorated the Poker Chip shale seal (Allan and Creaney,
149 1991). Peace River bitumen, which exhibits high levels of biodegradation, such as those from the
150 Bluesky Formation, are sourced by Exshaw-Banff source rock as these oils had a longer
151 residence time in the reservoir. On the other hand, deposits in the western part of Peace River
152 received more significant contributions from Gordondale source rock as these oils are higher in
153 sulfur content, less mature and have greater API gravity (Adams et al., 2013).

154

155 *2.2. Non-polar macromolecular fractions*

156 The experimental procedure used to obtain the non-polar macromolecular fraction (NPMF) from
157 Rozel Point and Jianghan oils was adapted from Adam et al. (1993). In brief, aliquots of the
158 crude oils (~100 mg) were loaded on a silica gel (6.0 g, 220-440 mesh) column. The samples
159 were washed with 50 mL of hexane and then eluted with 15 mL of hexane:diethyl ether (98:2),
160 which yielded the NPMF orange-colored band that was collected, evaporated to dryness and
161 submitted to (+) APPI FTICR-MS analysis. NPFM was not produced from western Canadian
162 oils indicating that the macromolecular fraction was not present in such mature and overmature
163 oils.

164 *2.3. Mass spectrometry*

165 Two primary ionization mechanisms are expected to occur in the APPI source: proton transfer,
166 which generates even electron protonated ions depending on the proton affinity of the analyte;
167 and the formation of an odd electron radical ion species (Raffaelli and Saba, 2003). Note that the
168 thermospray conditions in (+) APPI show negligible in-source thermal alteration of species
169 (Bagag et al., 2008).

170 Even though it was demonstrated in previous studies that (+) APPI FTICR-MS has some very
171 rudimentary quantitation capabilities (Oldenburg et al., 2014), due to unassessed differences in
172 ionization responses, intensities cannot be assumed to reflect the actual abundance of compounds
173 present in the sample. Thus, quantitative aspects of compounds and heteroatom classes are not
174 considered herein; only relative intensities are discussed, based on monoisotopic peak intensities
175 of assigned peaks (RMI, relative monoisotopic intensity). Also, structural discussions of detected
176 species are only speculative, since FTICR-MS data provide unambiguous molecular formulas but
177 cannot distinguish isomers.

178 Attempts to perform CID-FTICR-MS experiments in complex mixtures provide limited but
179 potentially useful information about the chemical structure of detected ions. Targeted peaks may
180 represent a dozen to potentially thousands of different isomers, as no prior chromatographic
181 clean-up was performed in these experiments. Therefore, the measured fragmentation patterns
182 represent the conflation of the fragmentation patterns of all the isomers with the targeted
183 molecular formula. The isolation window of the parent peak ($m/z \pm 0.4$) used in this study may
184 contain 10+ peaks. Thus, only peaks with the highest intensity within their $m/z \pm 0.4$ spectral

185 windows were selected for CID-FTICR-MS experiments to minimize interferences with non-
186 target peaks. This strategy limits the number of available targets, which have to be both
187 geochemically relevant and the dominant peak in its $m/z \pm 0.4$ spectral window. To further
188 reduce interferences with fragments originating from other parent peaks within the $m/z \pm 0.4$
189 windows, peaks with more carbon or sulfur atoms than the parent ion were ignored. Peaks
190 detected in CID-FTICR-MS experiments were assigned molecular formulae with assignment
191 errors < 100 ppb.

192 2.3.1. (+) APPI FTICR-MS method

193 The samples were analyzed using a 12 T Bruker Solarix FTICR-MS instrument. Whole oil and
194 fractions were diluted to 0.25 mg/mL in toluene and then infused into the ionization source using
195 a syringe pump set to deliver 200 $\mu\text{L}/\text{h}$. Atmospheric pressure photoionization in positive ion
196 mode via a krypton lamp at 10.6 eV was used to ionize the samples. Transfer capillary
197 temperature and nebulizer pressure were set to 350 $^{\circ}\text{C}$ and 1.0 bar, respectively. Reserpine
198 ($\text{C}_{33}\text{H}_{40}\text{N}_2\text{O}_9$) was added to the sample solution to assess internal calibration efficiency. The
199 instrument was tuned using a reference Athabasca whole bitumen sample. Ions ranging from m/z
200 150 to 1500 were isolated using a linear quadrupole and accumulated over 5 ms in the collision
201 cell before being transferred to the ICR cell. Spectra were collected in absorption mode, using an
202 algorithm proposed by Kilgour et al. (2013). Two hundred transients of 8 million points in the
203 time domain were collected and summed to improve the experimental signal/noise ratio (S/N).
204 FTICR-MS raw data were processed using the CaPA v.1.0 (Aphorist Inc.) software package.
205 Peaks with S/N higher than 4 were assigned based on highly accurate m/z measurements and on
206 stable isotopic distributions. Compositional boundaries, in terms of stoichiometry (element atom
207 content) for the fitting algorithm, were set to $\text{C}_{4-95}\text{H}_{0-200}\text{N}_{0-2}\text{S}_{0-8}\text{O}_{0-5}$ and double bond equivalent
208 (DBE) range, which is a measure of hydrogen deficiency due to double bonds and/or cyclic
209 structures, was limited to be between 0 and 60. The mass spectra were recalibrated using
210 homologous series present in the samples. Ragnarök v.2.0 (Aphorist Inc.) was used for data
211 manipulation and visualization. The resulting list of molecular formulae and intensities is
212 typically analyzed by summarizing the data based on three non-hierarchical layers: heteroatom
213 class, DBE and carbon number. Heteroatom class describes the amount of the heteroatoms
214 contained in a given molecular formula. The symbol “HC” is used for the class in which when no

215 heteroatom is present in the molecular formula, and classes detected as radical ions are
216 differentiated from protonated ones by a dot following the class symbol.

217 2.3.2. (+) APPI CID-FTICR-MS method

218 A set of peaks found in the oils was selected for CID experiments (Sections 4.3, 4.4 and 4.7).
219 Sulfur-bearing parent ions were isolated ($m/z \pm 0.4$) in the linear quadrupole and accumulated
220 over 0.3–1.0 s in the collision cell. The collision cell voltage was set to 12–20 V (compared to
221 the regular operating level of 5 V) to achieve collision-induced dissociation (CID) conditions,
222 causing the ions to undergo several low energy collisions with room temperature argon atoms.
223 Ions in the collision cell (parent and daughters) were then transferred to the ICR cell and the
224 CID-FTICR-MS spectra were collected in absorption mode after summing two hundred
225 transients of 4 million points in the time domain. The instrument was tuned to detect ions
226 ranging from m/z 150 to 1500, i.e. small molecular fragments ($< m/z$ 150) fell outside the
227 analytical window in these experiments. Bruker DataAnalysis v.4.4 and R v.3.4.0 software
228 packages were used for data processing and visualization.

229 3. Results

230 3.1. Non-polar macromolecular fraction

231 The NPMF fractions represented 18 and 22 wt% of the Jiangnan and Rozel Point oils,
232 respectively, and their (+) APPI mass spectra are shown in Figure 1.

233 3.2. (+) APPI FTICR-MS of oils and their fractions

234 In all spectra acquired in this study, the isotopic pattern of ^{13}C peaks indicated the presence of
235 singly charged species only. Averaged mass resolution ($m/\Delta m_{50\%}$) achieved for peaks detected
236 between m/z 397 and 403 was higher than 1,015,000 and higher than 560,000 for peaks between
237 m/z 797 and 803. Unique molecular formulas with absolute assignment error lower than 750 ppb
238 were assigned to 40,967 monoisotopic peaks present in the oil samples spectra (Fig. 1). In each
239 sample, peaks left unassigned represented less than 3% of total peak intensity and are not
240 discussed further. Heteroatom classes with fewer than 25 peaks were also excluded from the
241 discussion.

242 Fig. 2 shows the heteroatom class distribution of the investigated oils, as obtained by (+) APPI
243 FTICR-MS, where radical ion classes are differentiated from protonated ones by a dot following
244 the class symbol. The relative sulfur number distribution for the studied oils, as measured in
245 classes S_{1-6}^{\bullet} , is shown in Fig. 3A. Figs. 3B and 3C show the carbon number and DBE
246 distribution of all detected molecular formulae within the oil set. Fig. 4 highlights the DBE
247 distribution within the sulfur containing heteroatom classes, whereas Fig. 5 shows the carbon
248 number distribution of heteroatomic classes S_1 DBE 5 and 6 species. Fig. 6 shows the carbon
249 number distribution of heteroatomic classes S_1 DBE 1 and S_1^{\bullet} DBE 3. Fig. 7 shows the DBE
250 distribution of C_{40} species within class S_{1-5}^{\bullet} as found in Rozel Point and Jiangnan oils, whereas
251 Fig. 8 shows the heteroatomic class S_1^{\bullet} DBE 8 carbon distribution for the same oils.

252

253 3.3. (+) APPI CID-FTICR-MS

254 CID experiments were set to investigate the fragmentation patterns of a few selected parent ions
255 by measuring fragments with $m/z > 150$. Tables 2 and 3 show results of the CID-FTICR-MS
256 experiments used to investigate S_1 DBE 5 C_{27-30} and S_2^{\bullet} DBE 11 C_{56-57} species, respectively. The
257 target ion selection was based in the experiments described in Section 3.2 and will be further
258 discussed throughout Sections 4.3 – 4.7.

259 4. Discussion

260 Whereas the GC-MS, on which most of previous work is based, is limited to targeting specific
261 amenable compound isomers that may be present in very small concentrations, FTICR-MS data
262 represents conflation of individual compounds into compound classes, DBE and carbon
263 numbers, and are more representative of bulk properties. A direct comparison between previous
264 work that used a combination of desulfurization reactions and GC-MS analysis, and the current
265 study using FTICR-MS data is not possible given the different analytical capabilities of these
266 technologies, each with different limitations and advantages. Throughout this discussion section,
267 we refer to and use literature results based on GC-MS analysis to support interpretations based
268 on the FTICR-MS data. The ability to study molecules without the need of chemically modifying
269 the matrix via desulfurization reactions brings new observations worthy of discussion given their
270 impacts on the understanding of sulfurization reactions occurring in the geosphere.

271

272 *4.1. NPMF and sulfur-bridged (macro)molecules*

273 The current understanding of organic matter sulfurization mechanisms assumes the substitution
274 of functional groups in the precursor molecules by sulfur atoms, which then can rearrange to
275 more stable forms such as thiolanes and thiophenes, and/or act as bridges linking two precursor
276 molecules. As diagenesis progresses, such linkages would propagate and eventually generate a
277 Type II-S kerogen, which can produce oil molecules at a much lower level of thermal stress than
278 other kerogen types (Tegelaar et al., 1989; Vandenbroucke and Largeau, 2007).

279 Adam et al. (1993) reported a hexane soluble nonpolar macromolecular fraction (NPMF)
280 occurring in sulfur rich oils (including the RP oil), composed of highly aliphatic, high molecular
281 weight structures, possibly cross-linked by sulfur in a process similar to natural vulcanization.
282 Through a stepwise selective sulfur removal procedure, Richnow et al. (1993) proposed that in
283 RP oil, *n*-alkanes, hopanoids, steroids and phytanes are bound simultaneously via oxygen, sulfur
284 and aromatic units, based on the position of the functionality in the precursor molecules, creating
285 a macromolecular matrix. Alcohols released in the desulfurization of RP oil NPMF can be as
286 abundant as hydrocarbons and exhibited a similar carbon number distribution to hydrocarbons,
287 but the origin of most species could not be assigned (Jenisch-Anton et al., 1999). Efforts to
288 chemically or physically degrade such macromolecular organic matter were needed to
289 circumvent the inability of gas chromatography to analyze sulfur-rich macromolecules (Jenisch-
290 Anton et al., 1999). Investigating the RP oil, Adam et al. (1993) reported the presence of a
291 NPMF composed of cross-linked molecules containing multiple sulfur atoms, which represented
292 up to 32 wt% of the whole material (Adam et al., 1992; Jenisch-Anton et al., 1999). Significantly
293 different average molecular weight values of NPMF from RP oil were reported using various
294 techniques (Adam et al., 1993): 815 Da by vapor pressure osmometry; 3400 Da by light
295 scattering measurements and 1660 Da by size exclusion chromatography.

296 Despite a minor clean-up of non-S-bearing heteroatom classes that was applied during sample
297 preparation, FTICR-MS spectra of NPMF fractions of JH and RP oils analyzed in this study are
298 very similar to the respective whole oil spectra (compare panels A vs. B and C vs. D in Fig. 1),

299 indicating that the analyses of whole oils by (+) APPI are representative and capture the major S-
300 bearing constituents in those oils.

301 The spectra shown in Fig. 1 (A-G) revealed no detectable peaks above m/z 1100 in all the
302 analyzed oils and fractions. To further corroborate the absence of high molecular weight
303 compounds, we focused on the well-studied RP oil as representing very low maturity, high sulfur
304 oils. A (+) APPI FTICR mass spectrum was collected from m/z 1000–3000 (Fig. 1H), after
305 tuning the instrument with Agilent Low Concentration Tuning mix for enhanced sensitivity at
306 this m/z range. A group of peaks was detected in the RP oil at m/z 1000-1250 in Fig. 1H. Yet, no
307 evidence could be found for larger molecules ($m/z > 1500$). If the NPMF is formed by
308 consecutively adding building block units in a process similar to vulcanization, fragments at m/z
309 > 1250 would be expected since a continuum of species should exist (Eglinton et al., 1994). The
310 Jiangnan and Rozel Point oil NPMF spectra (Fig 1B and 1D, respectively) showed no indication
311 of ‘macromolecules’ in their composition. Notably, there is no evidence of in-source
312 fragmentation, consistent with previous work by Bagag et al. (2008) on the APPI analysis of
313 sensitive biomolecules, nor evidence of fragmentation during the ion transfer to the ICR cell. In
314 addition, the samples fully dissolved in toluene and no evidence of molecular aggregation was
315 seen visually or in the mass spectra. Hence, the (+) APPI FTICR-MS spectra obtained herein
316 represent a key piece in the understanding of the so called NPMF. The (+) APPI FTICR-MS
317 results clearly indicate that a macromolecular fraction >1250 Da is not present in the RP and JH
318 oils.

319 Intermolecular sulfur incorporation is inferred from the elevated relative intensity detected in the
320 C_{54-59} range within class $S\bullet_1$ DBE 8 (Fig. 8), tentatively interpreted to be the indication of two
321 C_{27-29} steroidal units bridged by one sulfur atom, each unit contributing 4 double bond
322 equivalents. Noteworthy, a similar pattern was seen in class $S\bullet_{2-3}$ DBE 9-11 C_{54-59} species, but
323 not in class $S\bullet_1$ DBE 7, which further suggests these species are indeed two bridged steroid units
324 (no DBE < 8) or eventually hopanoids for species with higher DBE. RP class $S_2\bullet$ DBE 11 C_{56-57}
325 compounds were selected as targets for CID experiments (Table 3) and results show detectable
326 daughter ions with the loss C_{28} and C_{29} with no sulfur and Δ DBE of 5.5, suggestive of an
327 intermolecular linkage.

328 Although they were inferred to exist, intermolecular sulfur incorporation reactions that yield high
329 molecular weight molecules are limited. Recent investigations have recognized the impact of
330 nanoaggregation on molecular weight measurements of other oil components, such as
331 asphaltenes (Zhang et al., 2013). Initially thought to be as high as several thousand kilodaltons,
332 the average molecular weights of non-aggregated asphaltene distributions are now recognized to
333 be around 750 Da (Mullins, 2010; Hosseini-Dastgerdi et al., 2015; Snowdon et al., 2016).
334 Similarly, an eventual aggregation of sulfurized species might be the reason for the
335 overestimation of NPMF molecular weights as measured via light scattering and size exclusion
336 chromatography in previous studies (Adam et al., 1993).

337 *4.2. Thermally immature vs. biodegraded sulfur rich oils*

338 Sinninghe Damsté and De Leeuw (1987) used gas chromatography coupled to mass spectrometry
339 (GC-MS), to identify not only sulfur containing isoprenoids (C₁₅ and C₂₀) in Rozel Point oils but
340 also a series of isoprenoid chains bonded to several organosulfur structures (e.g. thiophenes,
341 thiolanes, benzothiophenes). Further work led by the same authors expanded the scope of sulfur
342 compounds detected by GC-MS: C₃₀ and C₃₅ isoprenoid thiophenes, alkylthianes, isoprenoid
343 thiolanes, thiophene steranes, thiolane steranes, alkylbenzothiophenes, isoprenoid
344 benzothiophenes, isoprenoid bithiophenes and sulfur-containing hopanoids. In total, around one
345 thousand sulfur compounds were identified, with a molecular weight upper limit of 600 Da and a
346 maximum of two sulfur atoms in the structure (Sinninghe Damsté et al., 1987). Such species
347 were considered different from those sulfur compounds found in most oils since their structures
348 are closely related to biogeochemical precursors (Sinninghe Damsté et al., 1989). Lu et al. (2014)
349 investigated a set of Jianghan oils using (+) ESI FTICR-MS, after S-methylation reactions and
350 detected high levels of alkylcyclothioethers within the oils. The identified peaks ranged from
351 classes S₁₋₃, DBE 1-14, and C₁₀₋₃₅. Lu et al. (2014) also showed that compounds from class S₁
352 DBE 1-3 were detected with remarkably higher relative intensities and the predominance of C₂₀₋
353 ₂₁ compounds was attributed to the sulfurization of phytanic acids and phytols, and the overall
354 sulfurization mechanism in Jianghan oils was hypothesized to involve carboxylic acids and fatty
355 alcohols, which resulted in the formation of cyclic thioethers. Kohnen et al. (1993) identified
356 sulfur-bound steroid and phytane moieties in the Jianghan oil (also in Rozel Point oil),
357 suggesting that di- or polysulfide linkages are present in S-containing moieties and the position

358 of double bonds in the precursors controls the position of S-linkages. Liu et al. (2018) showed
359 that the ubiquitous sulfur ring structures present in the organic sulfur compounds in Jianghan
360 Basin oils likely originated from intramolecular sulfurization reactions.

361 The biogeochemistry of Peace River oils has been extensively studied in our group by Adams et
362 al. (2013) and Bennett et al. (2013). More recently, Oldenburg et al. (2017) have shown the
363 effects of different biodegradation levels in the chemical composition distribution in a Peace
364 River area, Bluesky Formation reservoir profile, as measured by FTICR-MS. Overall, molecules
365 with higher DBE and higher sulfur number were found to be more resistant to biodegradation.

366 The results of this study show that the mass spectra of Jianghan and Rozel Point oils are
367 remarkably different from the spectra of the Peace River oils (Fig. 1, A and C vs. E – G). The
368 smoother distribution of peaks detected in the Peace River oils, similar to a bell-shaped curve, is
369 a very common observation in fossil fuel analysis (Marshall and Rodgers, 2008) and relates to
370 the loss of carbon number preferences caused by the thermal cracking of kerogen during oil
371 generation. On the other hand, both Jianghan and Rozel Point oils show mass spectra that reflect
372 a dominant signature of precursor compounds (Fig. 3C), somewhat akin to the spectra observed
373 in recent marine sediments (Radović et al., 2016a), a consequence of the low thermal stress
374 levels experienced by these oils. Thus, both Rozel Point and Jianghan oil (+) APPI FTICR-MS
375 mass spectra are a good representation of the chemical species and mechanisms involved in early
376 diagenesis and related sulfur incorporation processes.

377 In these two oils, the predominance of heteroatom classes with one or more sulfur atoms (S_{1-6}) is
378 remarkable (Fig. 2), whereas classes HC, O_{1-2} , and NO_{0-1} are the only non-sulfur containing
379 heteroatom classes detected, accounting in total for less than 10% relative intensity in JH and RP
380 oils. The vast majority of compounds detected in these low maturity oils are sulfurized. In
381 addition, this is the first report of species with more than two sulfur atoms per molecule in Rozel
382 Point oils, while Liu et al. (2018) detected up to S_5 compounds in Jianghan oils.

383 The Peace River oils exhibit higher relative intensity for $S_{1\bullet}$ species, and a lower relative
384 intensity as the sulfur number increases up to $S_{4\bullet}$ (Fig. 3A). Both Rozel Point and Jianghan oils,
385 on the other hand, show maximum relative intensity at $S_{2\bullet}$ and extend up to $S_{6\bullet}$ and $S_{5\bullet}$,
386 respectively. Such differences may be associated with the depositional settings favoring organic

387 matter reaction with one or more sulfide species, coupled to the low thermal stress experienced
388 by the Jiangnan and Rozel Point oils. Therefore, the relative ratios of lesser and more extensively
389 sulfurized compound classes (e.g. S₁• vs. S₂•) is proposed as a proxy for source rock depositional
390 settings, although further testing with an extended sample set is required.

391 Peace River oils WC and BS show very similar chemical composition (Figs. 2 and 3). Despite
392 showing a similar carbon number distribution to its Peace River analogs in Fig. 3C, the Peace
393 River GR oil DBE distribution is significantly shifted towards lower DBE values (Fig. 3B),
394 while exhibiting a relative enrichment in S₂• species (Fig. 3A). Geochemical differences among
395 the Peace River oils studied herein have been observed before by Adams et al. (2013), including
396 a lower thermal maturity, lower biodegradation extent and higher sulfur content for oils primarily
397 sourced from the sulfur enriched Gordondale Formation.

398 *4.3. Sulfurized steroids and hopanoids (class S₁• DBE 5-6)*

399 In a study by Lu et al. (2013), C₂₈₋₃₀ DBE 5-7 steroids were detected in class S₁ of a sulfur rich
400 heavy oil in Jiaxian Sag, Bohai Bay Basin, China. Sulfurized steroidal structures were previously
401 reported in RP oil as part of the macromolecular matrix through a sulfur linkage located in ring
402 A or B (Adam et al., 1992). Thiophene and thiolane steranes have been identified in RP oil by
403 Sinninghe Damsté et al. (1987). Using selective cleavage of acyclic sulfide by superheated
404 methyl iodide, Schouten et al. (1993) described the release of C₂₇₋₃₀ steroid products from RP
405 polar and asphaltene fractions. Kohnen et al. (1993) suggested that multiple substrates prone to
406 sulfur incorporation reactions, such as Δ²-, Δ³-, Δ⁵-sterenes or Δ^{3,5}-steradienes, give rise to a
407 large variation in the positions and stereochemistry of sulfur incorporated steroids and that
408 timing of sulfur incorporation during diagenesis would significantly influence the resulting
409 products.

410 As can be observed in both in the overall (Fig. 3B) and class S₁• (Fig. 4A) DBE number
411 distribution plots, there is an elevated relative intensity of DBE 5-6 species in Rozel Point and
412 Jiangnan oils. C₂₇₋₃₀ species largely dominate the class S₁• DBE 5 carbon number distribution
413 (Fig. 5A). Since DBE 4 species do not show an elevated relative intensity for C₂₇₋₃₀ in any
414 compound classes (except for HC•), sulfurized steroid-like compounds likely include a
415 cyclothioether moiety to result in a DBE 5 value (Fig. 5A). Although present, C₂₇₋₃₀ species do

416 not exhibit an elevated relative intensity within heteroatom classes with more than 2 sulfur
417 atoms. This suggests that less functionalized molecular precursors such as steroids yield less
418 sulfurized diagenesis products, in contrast to precursors such as carotenoids that have multiple
419 reactive sites for sulfur incorporation (Section 4.5).

420 Results also suggest the occurrence of sulfurized homohopanes and norhopanes. The prominent
421 class $S_1 \bullet$ DBE 6 C_{35} peak (Fig. 5B) may represent the sulfurized analog of C_{35} homohopanes,
422 which are known to be the dominant triterpene in the Rozel Point oil. The elevated relative
423 intensity of class $S_1 \bullet$ DBE 6 C_{27-29} species in RP oil is remarkable (Fig. 5B) because there have
424 been no reports of S-bound norhopanes despite the occurrence of S-bound (homo)hopanes in
425 thermally immature sedimentary organic matter. Within the most plausible routes for the
426 production of sulfurized steroids in natural environments (Lu et al., 2013), none would yield S-
427 steroid species with DBE 6. To further investigate the forgoing interpretations, ions from class
428 $S_1 \bullet$ DBE 5 C_{27-30} were selected as targets for CID experiments (Table 2). In both RP and JH oils,
429 detected daughter ions represent the loss of C_{1-6} fragments with Δ DBE -0.5, i.e. aliphatic
430 moieties such as $-CH_3$, $-C_2H_5$, and no sulfur atoms, compatible with a side carbon chain
431 fragmentation from a D-ring or an AB-ring sulfurized steroid (Lu et al., 2013). The RP oil
432 targeted ions show, although with low relative intensity, $\Delta C_{0-2} \Delta S_1$ daughter ions, which indicate
433 that some class $S_1 \bullet$ DBE 5 C_{27-30} isomers might have thiol ($\Delta C_0 \Delta S_1$ fragments) or aliphatic
434 sulfide ($\Delta C_{1-2} \Delta S_1$) as functional groups, instead of cyclic sulfide. In such cases, the parent DBE
435 5 ion could represent, among other structures, either a 5-ring moiety (e.g., hopanoids) or a 4-ring
436 structure with one carbon-carbon double bond (e.g., sterenes). However, to the best of our
437 knowledge, no sterenes have been detected in RP oil, despite being thermally immature, thus the
438 occurrence of sulfurized norhopanes as class $S_1 \bullet$ DBE 6 C_{27-29} species is more likely.

439 *4.4. Class S_1 DBE 1 and class $S_1 \bullet$ DBE 3 species*

440 Previous work by Lu et al. (2014) indicated that the Jiangnan oils are relatively enriched in Class
441 S_1 DBE 1-3 species, and the same pattern was observed herein (Figs. 4A,B). The elevated
442 relative intensity of class S_1 DBE 1 species suggests an enrichment in thiolanes or thianes,
443 whereas elevated relative intensity of class $S_1 \bullet$ DBE 3 species suggests thiophenic structures.
444 There is an even/odd predominance spanning C_{16-40} class S_1 DBE 1 species in Jiangnan oil (Fig.

445 6A). Lu et al. (2014) reported the occurrence of odd/even predominance in the Jiangnan oil class
446 S₁, centered around the C₂₁ peak. Although not discussed in their paper, the methylation
447 reactions used by Lu et al. (2014) to facilitate (+) ESI detection of sulfur species added one
448 carbon to all the species detected in their study, and this might have caused the reversal of actual
449 even/odd carbon preference to observed odd/even preference in their data. In agreement with our
450 observations, Sinninghe Damsté et al. (1987) and Sheng et al. (1987) reported a series of C₁₀₋₃₂
451 alkylthiolanes and alkylthianes in Rozel Point and Jiangnan oils, respectively, which also
452 exhibited a strong even/odd preference. Here, the (+) APPI results show that the even/odd
453 preference extends much further, up to C₄₀ in the Jiangnan oil (Fig. 5A).

454 CID experiments were also performed to investigate the fragmentation patterns of class S₁ DBE
455 1 C_{26,28,30} species. However, the experiments failed to produce any detectable daughter ion at *m/z*
456 > 150. Typically, CID promotes the rupture of the weakest bonds in a molecule, therefore the
457 intermediate species generated after C-S bond cleavage in thiolanes/thianes might have
458 undergone extensive fragmentation. Based on the same reasoning, the presence of aliphatic thiols
459 or dialkylsulfides contributing to class S₁ DBE 1 species cannot be precluded, although the lack
460 of class S₁ DBE 0 species indicates such functional groups may be largely absent in the oil
461 matrices. Initially thought as intermediates in the formation of thiophenes, thiolanes have been
462 detected in sediment extract but their association with thiophene formation remains unclear
463 (Sinninghe Damsté et al., 1986). The (+) APPI FTICR-MS results support the hypothesis that (a)
464 these cyclic sulfides, formed during very early stages of diagenesis, likely never partake in the
465 kerogen formation (Brassell et al., 1988; Peng et al., 1998); and that (b) these cyclic sulfides may
466 be derived from the same functionalized precursors as the *n*-alkanes, such as *n*-alkanoic acids
467 and *n*-alkanols (Brassell et al., 1988; Liu et al., 2018). In this sense, reasonable follow up
468 research might incorporate a comprehensive assessment of sulfur-bearing species in sediment
469 extracts, occurring throughout a thermal maturity sequence, from recent immature sediments to
470 Type IIS kerogen at ~0.6 %Ro, to get insights on the dynamics of free molecules versus those
471 that are bound-then-released from kerogen.

472 Class S₁• DBE 3 species, which putatively represent thiophenic structures, show elevated relative
473 intensity at C_{20,30,35,40,45} in the Rozel Point oil, and at C_{20,24-26,28,30,40} in the Jiangnan oil (Fig. 5B).
474 FTICR-MS data alone are not capable of distinguishing isomers, but the elevated relative

475 intensity of multiple isoprene units ($x \cdot C_5$ units) may indicate the detection of an extended range
476 of isoprenoid thiophenes. To support this reasoning, isoprenoid thiophenes with < 35 carbons
477 have been identified by GC-MS in Rozel Point and Jiangnan oils, as reported by Sinnighe
478 Damsté et al. (1987) and Sheng et al. (1987), respectively.

479 4.5. *C₄₀ species*

480 The overall carbon distribution plot (Fig. 3C) shows an elevated relative intensity for C_{40} species
481 in both RP and JH oils. This observation alone suggests the presence of compounds bearing
482 much of the carbon skeleton from their biological precursor. Carotenoids are the likely biological
483 source for C_{40} species, since they are common constituents of living organisms with wide
484 taxonomic diversity, from archaea and cyanobacteria to higher plants and animals (Walter and
485 Strack, 2011). Recently, we used a similar (+) APPI FTICR-MS approach to characterize
486 carotenoids in recent marine sediments, in addition to various other lipid markers (Radović et al.,
487 2016a). Whereas C_{27-30} S-bearing steroids appear to be limited to classes S_{1-2} (see Section 4.3.),
488 C_{40} S-carotenoids are seen in classes with up to six sulfur atoms. Such findings indicate the clear
489 dependence of the degree of sulfur incorporation on the number of double bonds present in the
490 lipid undergoing diagenesis, supporting the observations and interpretations reported previously
491 with ESI FTICR-MS data (Liu et al., 2018). Amrani and Aizenshtat (2004) highlighted the
492 importance of carbonyl groups for the sulfurization processes in marine sediments, but the C_{40}
493 S-carotenoid species suggested here indicate that carbon-carbon double bonds may be driving the
494 sulfurization in such species. Carotenoids can offer multiple spots for sulfur incorporation,
495 contrary to steroids which have a limited number of double bonds and oxygenated groups. The
496 predominance of C_{40} in classes S_{5-6} is remarkable and reveals that, overall, most of the
497 sulfurized species may resemble the biological precursors in both Jiangnan and Rozel Point oils.
498 Further comparisons between sulfurization reactions in the laboratory versus natural settings
499 with various precursor molecules are required.

500 Typical carotenoid structures (e.g., lycopene, β -carotene, nostoxanthin) have 13 DBE units,
501 although higher DBE species are also possible (Walter and Strack, 2011). Carotenoid diagenesis
502 in different depositional settings can be quite complex (Watts and Maxwell, 1977; Repeta and
503 Gagosian, 1987), but the identification of several reduced carotenoids in recent sediments

504 suggests that hydrogenation occurs during diagenesis without structural or stereochemical
505 specificity (Hebting et al., 2006). Adam et al. (1993) not only suggested acyclic carotenes, β -
506 carotene and monocyclic carotenes as building blocks of Rozel Point NPMF, but also discussed
507 their implications to organic matter input determination. In the current study, similarly to Liu et
508 al. (2018), sulfurized carotenoids are present in the oils as free molecules instead of being
509 building blocks of NPMF, as evident by the absence of cross-linked carotenoid-derived NPMF
510 intermediates (i.e., no significant signal from C_{80} species was detected). In summary, the sulfur
511 incorporation in carotenoid precursors is not followed by an increase in the carbon number, thus
512 mostly original (biological) C_{40} species with varying DBE values were detected.

513 Fig. 7 shows the DBE distribution of classes $S_{1-5} \bullet C_{40}$ species. A significant intraclass pattern
514 observed in Fig. 7 is the elevated relative intensity for the DBE = $S + (2,4,6)$ species, which
515 might reflect the level of ‘thiophenization’. That is, DBE = $S + 0$ species would represent altered
516 aliphatic carotenoids (e.g. lycopene) where the sulfur atoms are incorporated as thiolane. DBE =
517 $S + 2$ species represent the altered carotenoids with 2 cyclic structures (e.g., beta-carotane) where
518 sulfur atoms are incorporated as thiolane. Note the higher relative intensity of DBE $S + 2$
519 compared to DBE $S + 0$ species, reflecting the higher abundance of cyclic carotenoid structure in
520 the deposited organic matter. Since the conversion of a thiolane into thiophenes involves a
521 Δ DBE +2 shift, species with DBE $S + 4$ and $S + 6$ may reflect the number of thiophenic
522 structures present in the compound. The transformation of thiolanes into thiophenes is likely
523 dependent on the thermal stress level experienced by the sample. The ratios of thiophene and
524 thiolane structures within the carotenoid classes require further investigation to determine its
525 potential usefulness as a marker for early diagenesis sulfurization.

526 4.6. The DBE number $\geq S_x$ rule

527 In Fig. 7, the minimum detected DBE number in each plot is equal to the number of sulfur atoms
528 of the corresponding class, i.e. DBE number $\geq S_x$, which indicates that, for C_{40} assignments, the
529 sulfur incorporation reactions are associated with a DBE increase. In fact, such a pattern is
530 observed in the dominant molecular formulae detected in this study, indicating that polysulfides,
531 aliphatic sulfides and thiols must be largely absent as stand-alone functional groups in the RP
532 and JH oils. The only exceptions to the rule are class $S_2 \bullet$ DBE 1 $C_{16,18,20,22,24,26}$ and class $S_3 \bullet$ DBE

533 2 C₃₇ species. The even/odd predominance in class S₁ DBE 1 was discussed previously. The
534 species in class S₂• DBE 1 C_{16,18,20,22,24,26} probably relates to those where an additional sulfur
535 atom is incorporated as a cyclic polysulfide or a thiol. The class S₃• DBE 2 C₃₇ species are
536 discussed in Section 4.7.

537 4.7. Sulfurized alkenones

538 Alkenones are long chain unsaturated ketones produced by some phytoplankton species and
539 typically used as paleoenvironmental proxies (Volkman et al., 1980; Brassell et al., 1986). An
540 elevated relative intensity was observed for classes S₁₋₃• DBE 1-3 C₃₇ species in Rozel Point oil,
541 suggesting that alkenones have also undergone sulfurization during diagenesis. Although it is
542 tempting to try and estimate the ratios of parent C_{37:2} and C_{37:3} (e.g., $U_{37}^{K'}$) based on their
543 sulfurized analogs, this exercise would be highly speculative at this point. Also, the occurrence
544 of class S₃• DBE 2 C₃₇ species as an exception to the DBE number $\geq S_x$ rule, indicates that in
545 some cases thiol or polysulfides might be present. The peak representing class S₂• DBE 2 C₃₇ in
546 Rozel Point oil was selected for CID experiments, and similarly to class S₁ DBE 1 peaks, no
547 daughter ions could be detected, except for an ion representing the loss of a -SH group, although
548 results do not clarify where this functional group occurs in the alkenone molecules. This result
549 suggests that the C₃₇ assignments investigated herein may display long alkyl chains that, in turn,
550 produce daughter ions, which undergo further fragmentation outside of our analytical window at
551 $m/z > 150$.

552 4.8. Implications for atmospheric carbon dioxide removal

553 To meet the challenges of climate predictions at the end of the 21st century (IPCC, 2018),
554 technologies will be needed that remove carbon dioxide from the atmosphere and sequester it for
555 geological time frames. Such processes, termed carbon dioxide removal (CDR) are increasingly
556 attracting both research and commercial investment interest. Only a few CDR technologies are
557 currently being commercially developed (e.g., bio-energy with carbon capture and storage, or
558 direct air capture), but are still far below the scale needed for globally significant CO₂ reduction,
559 which would require achieving carbon drawdowns in the order of 100 to 1000 Gt/y by the year
560 2100 (IPCC, 2018). A better understanding of natural mechanisms for long term preservation of
561 organic molecules can provide guidelines for the development of technologies that could

562 leverage those mechanisms to artificially enhance the removal of carbon in the geosphere in the
563 form of engineered organic species. We term these species “alternative vectors for carbon
564 storage”, or AVECS for short, and are currently investigating several routes to produce AVECS
565 from organic precursors, including pathways involving sulfurization of organic molecules (Yim,
566 2019). In one of the concepts under investigation, waste biomass and sulfur, which is a by-
567 product of sour petroleum production, are the feedstocks for a long-term subsurface preservation
568 of organic carbon in the geosphere. To that end, the study reported herein provided multiple
569 indications relevant to AVECS technology development issues such as:

- 570 • Natural precursors such as pigments or lipids can incorporate sulfur as free molecules;
- 571 • Double bonds are the key molecular site for sulfur incorporation;
- 572 • Given the fact that double bonds are quite reactive, the sulfur incorporation must be
573 taking place early in the diagenesis, at low pressures and temperatures, which favorably
574 implies fewer energy needs for AVECS production;
- 575 • Cyclization is the favored sulfur incorporation mechanism producing thiolane structures,
576 which can then be aromatized to thiophenes during diagenesis.

577 Some of these processes can be modelled and reproduced in the laboratory (Yim, 2019), however
578 much more research effort will be needed to move these types of reactions from bench-scale to
579 sustainable pilot or large-scale deployments; notwithstanding, biogeochemical processes
580 occurring in geosphere, such as those inferred from the current study of sulfur rich oils, can
581 provide valuable insights and lessons that can direct this type of research.

582 **5. Conclusions**

583 The present study leveraged the expanded analytical window offered by novel analytical
584 technologies to improve the current geochemical understanding of diagenetic processes
585 involving sulfur and the resultant occurrence of sulfur rich oils. A set of five sulfur rich oils was
586 analyzed via (+) APPI FTICR-MS and CID-FTICR-MS. From these results, we conclude:

587 The absence of peaks at $m/z > 1250$ in Rozel Point and Jiangnan oils, in their non-polar
588 macromolecular fractions does not support the concept of a macromolecular network linked by
589 (poly)sulfide bridges that was proposed to explain the occurrence of non-GC-MS amenable sulfur

590 compounds in immature, sulfur rich oils. Instead, it appears that the presence of highly sulfurized
591 lipids (up to six sulfur atoms) as free molecules are the result of sulfurization with the degree of
592 sulfur incorporation dependent on the abundance of functional groups (double bonds mainly) in
593 the precursor molecules.

594 Source rock thermal stress levels were found to be a key factor differentiating the composition of
595 biodegraded, sulfur-rich Peace River oils, compared to thermally immature, sulfur-rich Rozel
596 Point and Jiangnan oils. Whereas Peace River oils do not show biological precursor skeleton
597 signatures, Rozel Point and Jiangnan oils show the presence of S-containing steroids, hopanoids,
598 carotenoids and alkenones among others.

599 (+) APPI CID-FTICR-MS experiments indicated only a minor occurrence of thiol functional
600 groups and sulfide bridges linking two precursor molecules. The compositional patterns observed
601 in sulfurized C₄₀ assignments, representing S-bound carotenoids, supports a mechanism of sulfur
602 incorporation involving the formation of thiolane structures (and to a lesser extent, thianes),
603 which can then be aromatized to thiophenes during diagenesis. CID-FTICR-MS experiments
604 were also helpful to prove the occurrence of intermolecular sulfur incorporation reactions as a
605 secondary mechanism, the aliphatic nature of class S₁ DBE 1 species and sulfurized alkenones,
606 the preference of cycle-forming sulfur incorporation in steroid structures, and the occurrence of
607 thiol/thioether groups, although these were seen only to a minor extent.

608 The analytical approach presented here can be improved and expanded using analytical
609 workflows that include chromatographic separation to better characterize individual molecular
610 species, as well as by studying other sample types such as source rocks and recent sediments to
611 better constrain the progression of diagenetic processes involving sulfur incorporation.

612 The occurrence of S-bound compounds in oils or sediment extracts, as detected via FTICR-MS,
613 can be used as proxies for depositional environment, organic matter input, sulfurization and other
614 diagenetic processes. Understanding the relationship between the precursor and sulfurized
615 products, as well as the mechanisms for such reactions, is key to the design of geoengineering
616 solutions to enhance carbon preservation in the geosphere by promoting sulfurization of biomass.

617

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621 **Notes**

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631 contributed to writing. FTICR-MS method optimization and analyses were carried out by MB,
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842 **Table 1.** Sample codes and geochemical overview.

Code	Oil	Location	Sulfur (wt %)	Organic matter source	Depositional environment	Thermal maturity	Biodegradation	References
RP	Rozel Point	Utah, USA	9.4	Marine - lacustrine	hypersaline, lacustrine	Immature	Moderate	(Meissner et al., 1984; ten Haven et al., 1988; Sinninghe Damsté et al., 1989)
JH	Jiangnan	Eastern China	5.6	Marine - lacustrine, with terrigenous inputs	hypersaline, lacustrine	Immature	Moderate	(Philp and Zhaoan, 1987; Carroll and Bohacs, 2001; Hou et al., 2017)
GR	Peace River	NW-Central Alberta, Canada	5.5	Marine - lacustrine, with terrigenous inputs	Marine - lacustrine	Mature - overmature	High	(Riediger, 1994; Adams et al., 2012; Adams et al., 2013; Bennett et al., 2013)
BS	Peace River	NW-Central Alberta, Canada	4.1	Marine - lacustrine	Marine - lacustrine	Mature - overmature	Very High	
WC	Peace River	NW-Central Alberta, Canada	4.2	Marine - lacustrine	Marine - lacustrine	Mature - overmature	High	

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Table 2. (+) APPI CID-FTICR-MS results of class S₁ DBE 5 C₂₇₋₃₀ species. Relative intensities are normalized to the most intense daughter ion as the target peak is considered in the relative intensity calculations. Fragments with carbon or sulfur number higher than the parent peak, as well as those with relative intensity <10%, are not shown. Note: *lost of an even number of hydrogens, **fragment is likely not originated from the parent ion.

Parent	Oil	Fragment	Lost	Relative intensity (%)	Parent	Oil	Fragment	Lost	Relative intensity (%)	
C ₂₇ H ₄₆ S	Jianghan	C ₂₆ H ₄₃ S	-CH ₃	100.0	C ₂₉ H ₅₀ S	Jianghan	C ₂₈ H ₄₇ S	-CH ₃	100.0	
		C ₂₂ H ₃₅ S	-C ₅ H ₁₁	99.3			C ₂₃ H ₃₇ S	-C ₆ H ₁₃	83.5	
		C ₂₄ H ₃₉ S	-C ₃ H ₇	40.7			C ₂₂ H ₃₅ S	-C ₇ H ₁₅	33.4	
		C ₂₃ H ₃₇ S	-C ₄ H ₉	32.2			C ₂₁ H ₃₃ S	-C ₈ H ₁₇	22.9	
	Rozel Point	C ₂₆ H ₄₃ S	-CH ₃	100.0			C ₂₆ H ₄₃ S	-C ₃ H ₇	22.5	
		C ₂₂ H ₃₅ S	-C ₅ H ₁₁	84.9			C ₂₄ H ₃₉ S	-C ₅ H ₁₁	22.3	
		C ₂₄ H ₃₉ S	-C ₃ H ₇	19.5			C ₂₇ H ₄₅ S	-C ₂ H ₅	18.2	
		C ₂₃ H ₃₇ S	-C ₄ H ₉	16.2			Rozel Point	C ₂₈ H ₄₇ S	-CH ₃	100.0
		C ₂₇ H ₄₅	-SH	15.5		C ₂₃ H ₃₇ S		-C ₆ H ₁₃	53.7	
		C ₁₉ H ₂₉ S	-C ₈ H ₁₇	13.8		C ₂₂ H ₃₅ S		-C ₇ H ₁₅	26.6	
	C ₂₆ H ₄₁	-CH ₅ S	12.9	C ₂₆ H ₄₃ S		-C ₃ H ₇		25.5		
	C ₂₁ H ₃₄ S	-C ₆ H ₁₂ *	12.3	C ₂₈ H ₄₅		-CH ₅ S		19.2		
	C ₁₇ H ₂₆ S	-C ₁₀ H ₂₀ *	10.2	C ₂₄ H ₃₉ S		-C ₅ H ₁₁		16.2		
	C ₂₈ H ₄₈ S	Jianghan	C ₂₇ H ₄₅ S	-CH ₃		100.0	C ₂₉ H ₄₉	-SH	13.9	
C ₂₃ H ₃₇ S			-C ₄ H ₉	89.3	C ₂₇ H ₄₅ S	-C ₂ H ₅	13.5			
C ₂₂ H ₃₅ S			-C ₅ H ₁₁	76.1	C ₂₁ H ₃₃ S	-C ₈ H ₁₇	12.5			
C ₁₂ H ₁₅ S			-C ₁₆ H ₃₃	38.3	C ₃₀ H ₅₂ S	Jianghan	C ₂₉ H ₄₉ S	-CH ₃	100.0	
C ₂₅ H ₄₁ S			-C ₃ H ₇	31.2			C ₂₅ H ₄₁ S	-C ₅ H ₁₁	26.0	
C ₂₁ H ₃₃ S			-C ₆ H ₁₃	27.7			C ₂₄ H ₃₉ S	-C ₆ H ₁₃	14.0	
C ₂₄ H ₃₉ S			-C ₄ H ₉	26.5			C ₂₈ H ₄₇ S	-C ₂ H ₅	11.3	
Rozel Point		C ₂₇ H ₄₅ S	-CH ₃	100.0			C ₂₇ H ₄₅ S	-C ₃ H ₇	10.7	
		C ₂₂ H ₃₅ S	-C ₆ H ₁₃	44.4			Rozel Point	C ₂₉ H ₄₉ S	-CH ₃	100.0
		C ₂₃ H ₃₇ S	-C ₅ H ₁₁	41.7				C ₂₉ H ₄₇ S	-CH ₅	22.3
		C ₂₅ H ₄₁ S	-C ₃ H ₇	19.7	C ₂₄ H ₃₉ S	-C ₆ H ₁₃		21.4		
		C ₂₁ H ₃₃ S	-C ₇ H ₁₅	17.2	C ₂₃ H ₃₇ S	-C ₇ H ₁₅		21.1		
		C ₂₈ H ₄₇	-SH	16.1	C ₂₈ H ₄₇ S	-C ₂ H ₅		18.0		
		C ₂₇ H ₄₃	-CH ₅ S	14.0	C ₁₇ H ₂₅ S	-C ₁₃ H ₂₇		16.2		
C ₁₉ H ₂₉ S	-C ₉ H ₁₉	13.4	C ₂₉ H ₄₇	-CH ₅ S	16.2					
				C ₂₇ H ₄₅ S	-C ₃ H ₇	15.7				
				C ₂₈ H ₄₇	-C ₂ H ₅ S	14.1				
				C ₂₅ H ₄₁ S	-C ₅ H ₁₁	13.4				
				C ₂₈ H ₄₃ S	-C ₂ H ₉ **	12.2				
				C ₂₀ H ₃₁ S	-C ₁₀ H ₂₁	10.0				

Table 3. (+) APPI CID-FTICR-MS results of class S₂• DBE 11 C₅₆₋₅₇ species in the Rozel Point oil. Relative intensities are normalized to the most intense daughter ion as the target peak is considered in the relative intensity calculations. Fragments with carbon or sulfur number higher than the parent peak, as well as those with relative intensity <10%, are not shown.

Parent	Fragment	Lost	Relative intensity (%)
C ₅₆ H ₉₂ S ₂	C ₅₀ H ₇₉ S ₂	-C ₆ H ₁₃	100.0
	C ₅₁ H ₈₁ S ₂	-C ₅ H ₁₁	60.5
	C ₅₅ H ₈₉ S ₂	-CH ₃	14.3
	C ₂₇ H ₄₅	-C ₂₉ H ₄₇	13.4
	C ₅₂ H ₈₃ S ₂	-C ₄ H ₉	10.5
C ₅₇ H ₉₄ S ₂	C ₅₁ H ₈₁ S ₂	-C ₆ H ₁₃	100.0
	C ₅₂ H ₈₃ S ₂	-C ₅ H ₁₁	100.0
	C ₅₆ H ₉₁ S ₂	-CH ₃	18.5
	C ₂₉ H ₄₉	-C ₂₈ H ₄₅ S ₂	16.0
	C ₂₈ H ₄₇	-C ₂₉ H ₄₇ S ₂	14.2

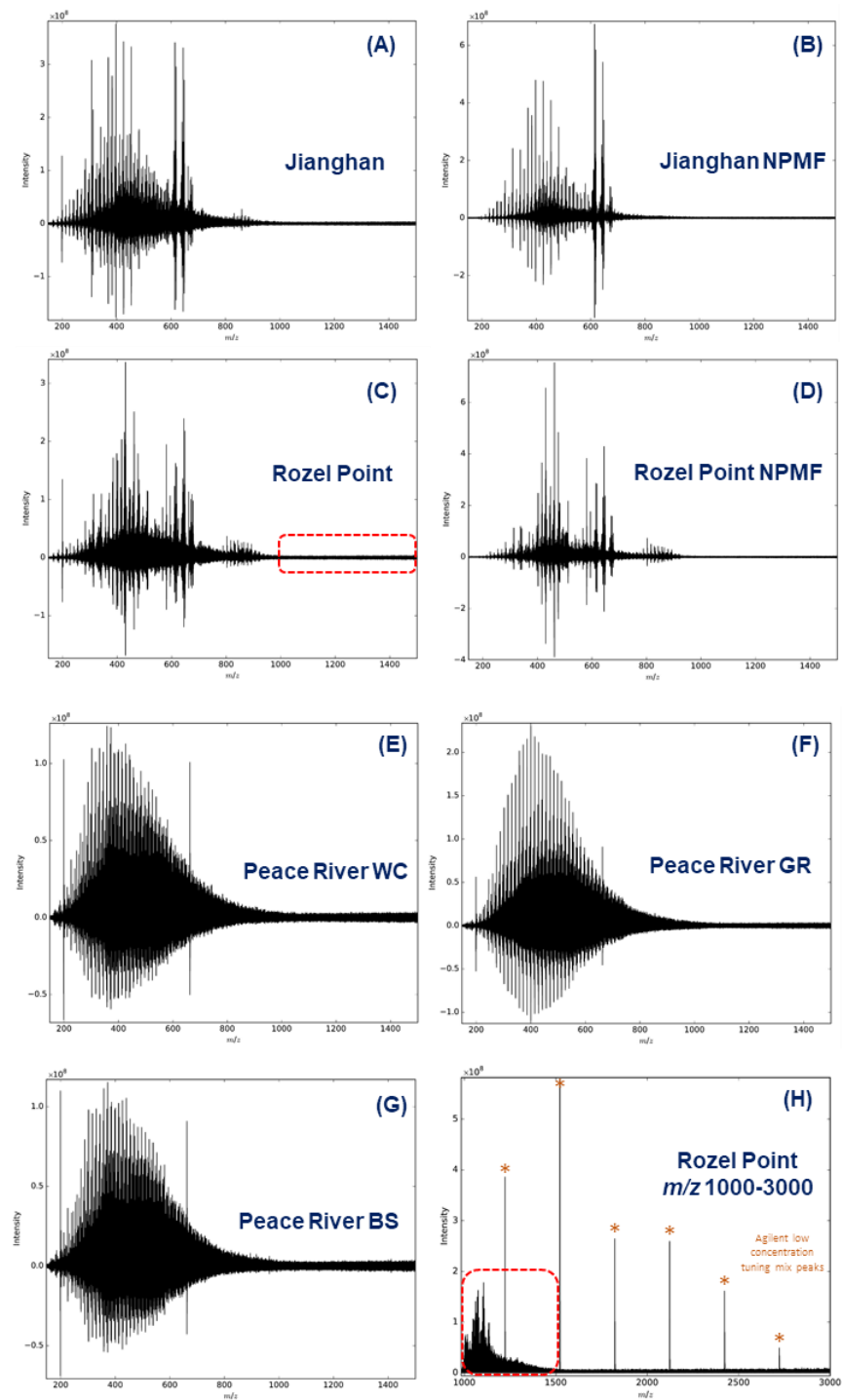


Figure 1. (+) APPI FTICR absorption-mode mass spectrum of (A-B) Jianghan oil and NPMF, (C-D) Rozel Point oil and NPMF, (E-G) Peace River oils (see Table 1 for sample codes). Plot H shows the Rozel Point oil APPI-P FTICR magnitude-mode mass spectrum, acquired from m/z 1000 – 3000, 100 scans, ion accumulation time increased by a 100-fold (500 ms), where the sample solution was doped with 2.5 μL of the Agilent Low Concentration Tuning Mix.

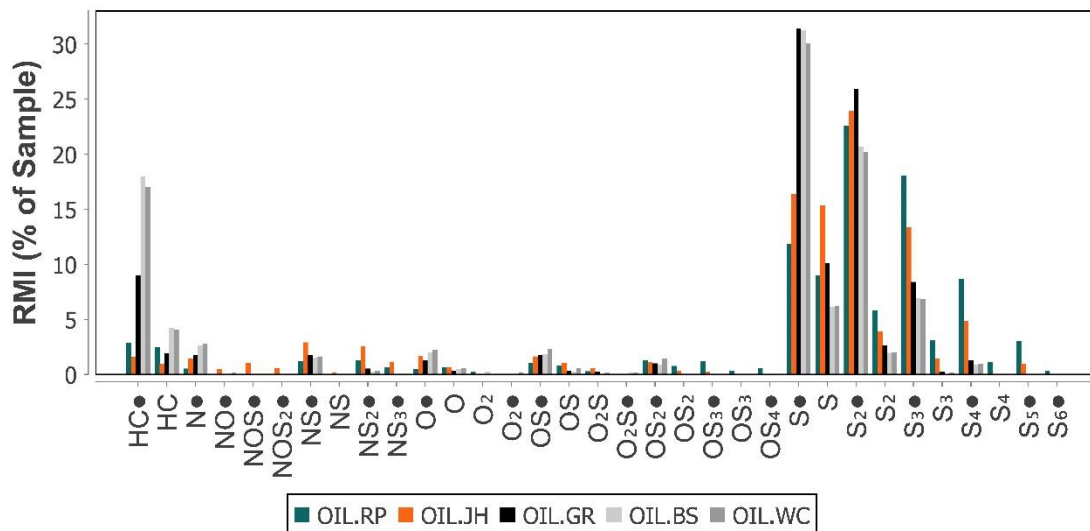


Figure 2. (+) APPI FTICR-MS compound class distribution of the oil samples analyzed in this study. Dots following class labels refer to radical ions (odd electron ion) while the absence of a dot refers to protonated ions (even electron ion). The relative monoisotopic intensity (RMI, % of Sample) is calculated as the fraction of the total monoisotopic intensities of a compound class, normalized to the sum of monoisotopic intensities of all compound classes in the sample. See Table 1 for sample abbreviations.

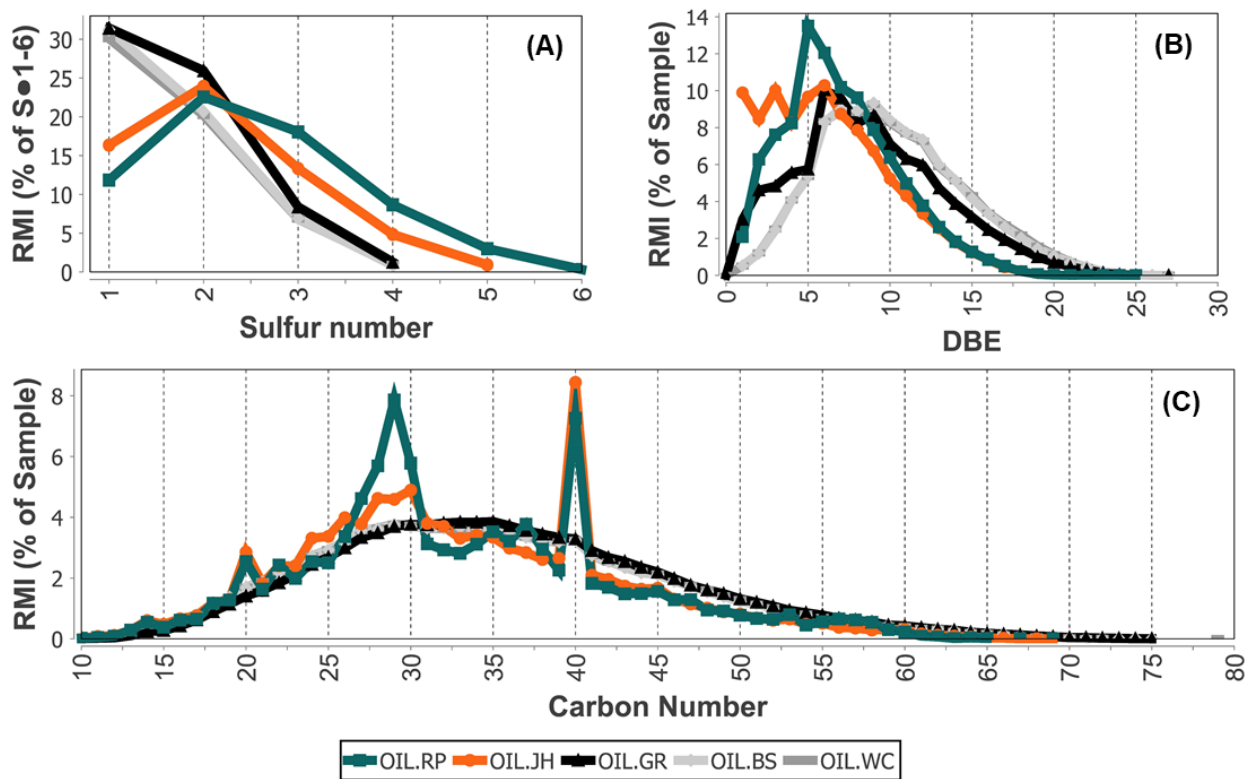


Figure 3. (+) APPI FTICR-MS parameter distributions in the oil samples analyzed in this study. (A) Sulfur number distribution, based on radical heteroatomic classes $S\bullet_{1-6}$ peaks; (B) DBE number distribution of all assigned peaks; (C) Carbon number distribution of all assigned peaks. See Table 1 for sample abbreviations.

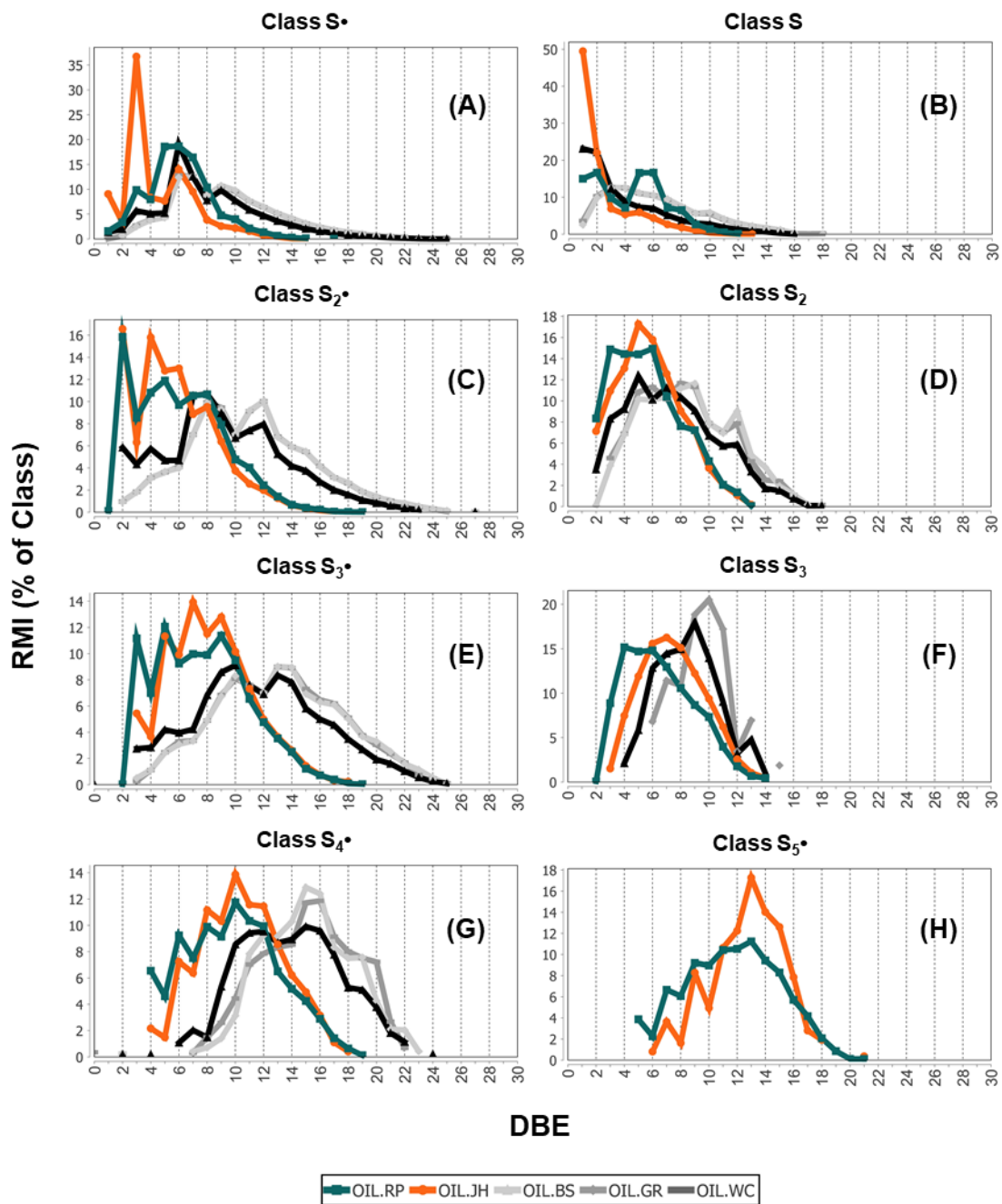


Figure 4. (+) APPI FTICR-MS heteroatomic classes S₁₋₅• and S₁₋₃ DBE number distributions in the oil sample set. Dots following class labels refer to radical ions (odd electron ion) while the absence of a dot refers to protonated ions (even electron ion). See Table 1 for sample abbreviations.

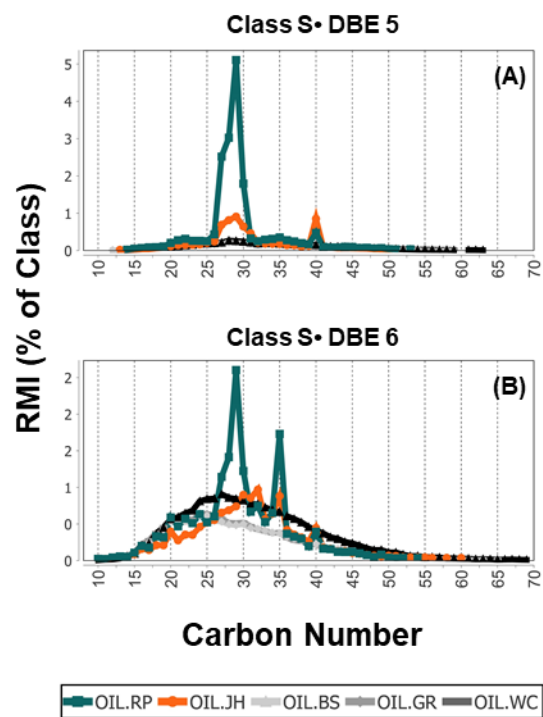


Figure 5. (+) APPI FTICR-MS carbon number distribution of heteroatomic classes S₁• DBE 5 (A) and 6 (B). Dots following class labels refer to radical ions (odd electron ion. See Table 1 for sample abbreviations). Sulfurized steroid-like C₂₇₋₃₀ species largely dominate the class S₁• DBE 5 whereas sulfurized (homo- and nor-) hopanoid-like species can be inferred from class S₁• DBE 6 carbon number distribution.

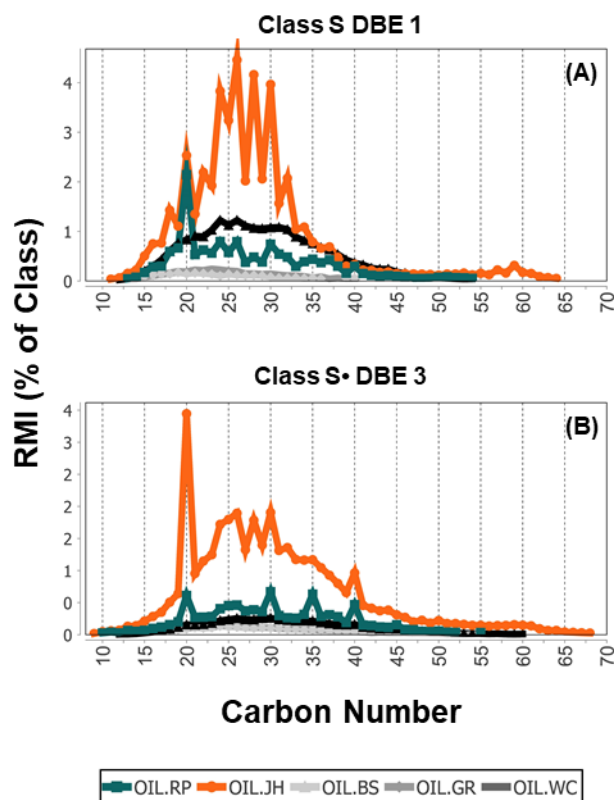


Figure 6. APPI-P FTICR-MS carbon number distribution of heteroatomic classes (A) S_1 DBE 1 and (B) $S_1\bullet$ DBE 3. Dots following class labels refer to radical ions (odd electron ion) while the absence of a dot refers to protonated ions (even electron ion). See Table 1 for sample abbreviations. The even/odd predominance spanning C_{16-40} class S_1 DBE 1 species in Jiangnan is putatively associated with alkylthiolanes and alkylthianes. The elevated relative intensity of multiple isoprene units ($x \cdot C_5$ units) in class $S_1\bullet$ DBE 3 species suggests an extended range of isoprenoid thiophenes in both oils but mainly in RP.

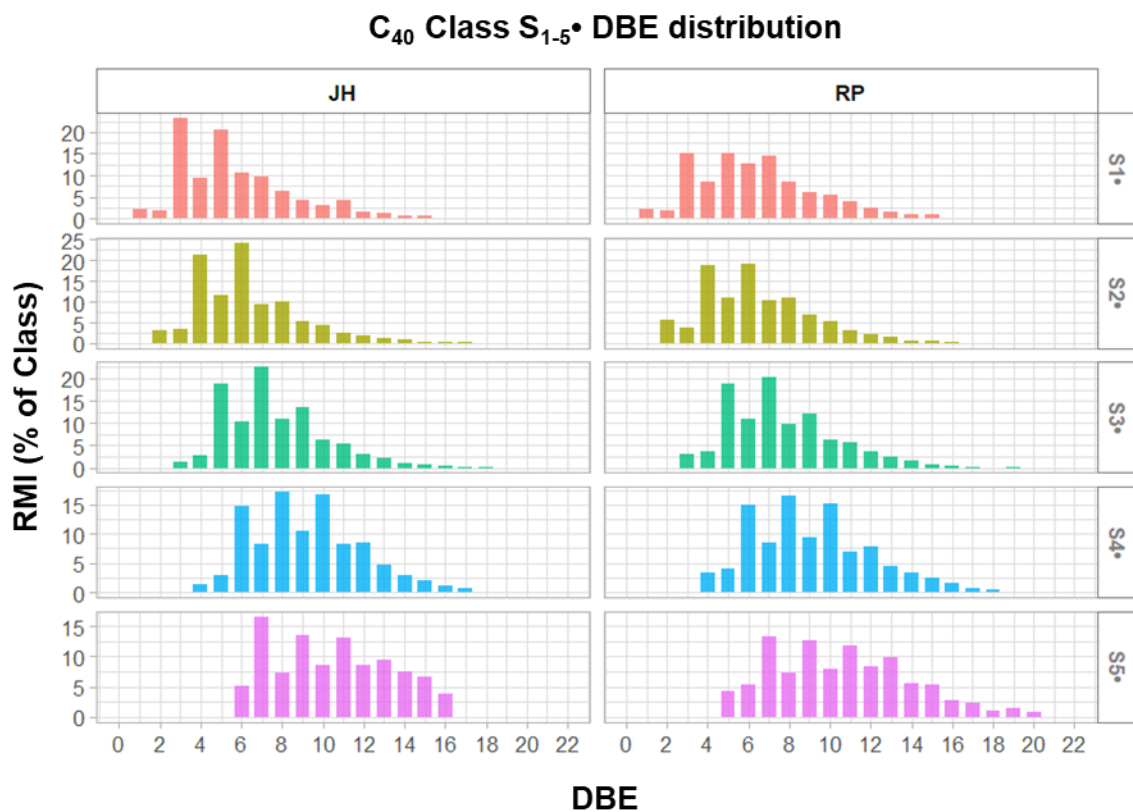


Figure 7. DBE distribution of class S₁₋₅• C₄₀ species in Rozel Point and Jiangnan oils, as detected by APPI-P FTICR-MS. An intraclass DBE = S + (2,4,6) elevated relative intensity pattern is observed and putatively associated with the number of thiophenic units in the molecules.

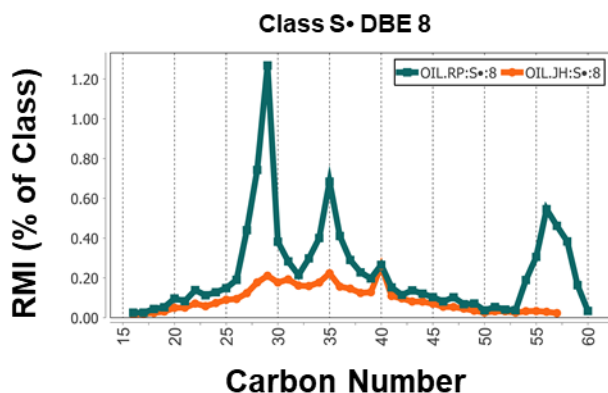


Figure 8. APPI-P FTICR-MS carbon number distribution of heteroatomic class S•₁ DBE 8, in Jiangnan (JH) and Rozel Point (RP) oils. Dots following class labels refer to radical ions (odd electron ion). The increased elevated RMI relative intensity detected in the C₅₄₋₅₉ range within class S•₁ DBE 8 is interpreted as two C₂₇₋₂₉ steroidal units bridged by one sulfur atom.