

1 **Mechanistic insights into sulfur rich oil formation, relevant to geological carbon storage**  
2 **routes. A study using APPI-P FTICR-MS(/MS) analysis**

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7 **Abstract**

8 Sulfur incorporation into sedimentary organic matter has a key role in carbon preservation in the  
9 geosphere. Such processes can inform strategies for human timescale carbon storage to mitigate  
10 climate change impacts and thus more detailed knowledge of sulfur incorporation into biomass  
11 species is needed. Until recently, detailed chemical characterization of sulfurized organic matter  
12 was only possible by analyzing individual building blocks obtained after desulfurization  
13 reactions. In this study, Fourier transform ion cyclotron resonance mass spectrometry (FTICR-  
14 MS), with atmospheric pressure photoionization in positive ion mode (APPI-P) was used to  
15 investigate the chemical composition of sulfur rich crude oils and to obtain mechanistic insights  
16 into the sulfur incorporation reactions happening during early diagenesis. Contrary to the current  
17 body of knowledge, APPI-P FTICR-MS data show that sulfurized lipids (with up to 6 sulfur  
18 atoms) occur as free molecules in these oils, rather than within a macromolecular network linked  
19 by (poly)sulfide bridges. In contrast to Peace River (Canada) oils, the thermally immature Rozel  
20 Point (USA) and Jiangnan Basin (China) oils show a carbon number preference in sulfurized  
21 species resembling biogenic precursor molecules, which highlights the importance of S-bound  
22 molecules as geochemical proxies for early diagenetic processes. This study indicates that sulfur  
23 incorporation reactions involve the formation of S-cyclic structures in which the double bond  
24 equivalent is  $\geq$  the number of S atoms, which can then aromatize to thiophenes depending on the  
25 thermal stress level they experience. FTICR-MS/MS experiments suggest the occurrence of  
26 intermolecular sulfur incorporation reactions, but only as a mechanism secondary to  
27 intramolecular sulfur addition. The FTICR-MS/MS fragmentation studies also indicate the  
28 aliphatic nature of class S<sub>1</sub> species and sulfurized alkenones, the preference of cycle-forming  
29 sulfur incorporation in steroid structures, and the occurrence, to a minor extent, of thiol/thioether

30 groups. Knowledge of organic sulfur molecule formation informs routes for carbon dioxide  
31 removal technologies which could sequester carbon in the geosphere and/or hydrosphere in the  
32 form of recalcitrant organic species.

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## 34        **1. Introduction**

35        Sulfur incorporation into sedimentary organic matter has a key role in carbon preservation in the  
36        geosphere, occurring on short geological timescales, not only on Earth but also possibly on Mars  
37        (Werne et al., 2004; Summons et al., 2011). The understanding of mechanisms and geochemical  
38        implications of sulfur incorporation reactions expanded greatly during the late 1980s and early  
39        1990s (Orr and Sinninghe Damsté, 1990; Aizenshtat et al., 1995; Amrani, 2014), and given their  
40        relevant role in enabling an important carbon sink, there are still many active areas of research in  
41        the field. For example, there currently exists a heightened research interest to further understand  
42        carbon preservation mechanisms via sulfur incorporation (Raven et al., 2016; Pohlabein et al.,  
43        2017) in the context of the global change and the need to keep the global temperature increase to  
44        <2 °C by the end of this century (IPCC, 2018).

45        Large scales of carbon removal are not unusual in natural systems. The global ocean has pools of  
46        organic carbon found in dissolved and particulate form in the water column, or deposited in the  
47        deep-sea sediments which are of similar magnitude to atmospheric carbon reservoirs (Houghton,  
48        2007; Hansell et al., 2009; Jiao et al., 2010). Therefore, a better understanding of natural  
49        mechanisms for long term preservation of organic molecules in the geosphere can provide  
50        guidelines for the development of technologies that artificially enhance the removal of carbon  
51        from the atmosphere to ocean or subsurface reservoirs, in the form of biologically refractory  
52        organic species analogous to those already present naturally in the Earth.

53        During the early diagenesis, sulfurized organic compounds, commonly occurring in anoxic  
54        environments that are rich in reduced sulfur, would typically resemble their biological  
55        precursors. As the diagenesis progresses, extensive sulfurization reactions enable multi-point,  
56        intermolecular cross-linkage, eventually generating the macromolecular structures which are the  
57        building blocks of sulfur rich kerogen (type II-S) (Tegelaar et al., 1989; Vandenbroucke and  
58        Largeau, 2007). Sulfur rich macromolecular structures were also detected in thermally immature,  
59        sulfur rich oils (Sinninghe Damsté et al., 1987; Sinninghe Damsté et al., 1989; Orr and Sinninghe  
60        Damsté, 1990). However, the detailed chemical characterization of such compounds was only  
61        possible by analyzing individual building blocks after desulfurization reactions (Adam et al.,  
62        1992; Adam et al., 1993; Richnow et al., 1993). Since then, analytical technologies have greatly  
63        advanced. Ultrahigh resolution mass spectrometry, in particular the FTICR-MS (Fourier

64 transform ion cyclotron resonance mass spectrometry) technology, has enabled the identification  
65 of several thousand peaks in complex organic mixtures, revolutionizing the understanding of the  
66 petroleome (Marshall and Rodgers, 2008). In the field of organic geochemistry, such technical  
67 advances were translated into an unprecedented capability to probe the high molecular weight,  
68 polar compounds at a high level of detail, expanding the way biodegradation, thermal stress and  
69 other biogeochemical processes are understood (Oldenburg et al., 2014; Radović et al., 2015;  
70 Radović et al., 2016b; Oldenburg et al., 2017).

71 Investigations focused on the geochemical significance of sulfur species have also greatly  
72 benefited from ultrahigh resolution mass spectrometry applications. Hughey et al. (2004)  
73 analyzed two same sourced Smackover Formation oils of different levels of thermal maturity  
74 using ESI-N (electrospray ionization in negative ion mode) FTICR-MS, suggesting the removal  
75 of sulfur and oxygen compounds was promoted by thermal maturation. Similarly, Oldenburg et  
76 al. (2014) reported that the relative apparent abundances of all heteroatom-containing compound  
77 classes (nitrogen, oxygen, sulfur and mixed heteroatom species) detected in that study decreased  
78 systematically with increasing oil maturation levels. Walters et al. (2015) used both APPI-P  
79 (atmospheric pressure photoionization in positive-ion mode) and ESI-N FTICR-MS to probe  
80 sulfur compounds and oxygenated analogs present in samples from the Smackover Formation,  
81 revealing intermediates (classes  $O_x$  and  $SO_x$ ) of thermochemical sulfate reduction (TSR) redox  
82 reactions. A similar approach detected the formation of TSR induced proto-solid bitumen,  
83 including highly condensed polynuclear aromatic and naphthoaromatic species with up to  
84 three sulfur atoms (Walters et al., 2011). Lu et al. (2013) used ESI-P (electrospray ionization in  
85 positive-ion mode) FTICR-MS to investigate sulfur rich heavy oils from Bohai Bay Basin,  
86 reporting a wide range of sulfurized sedimentary steroids in addition to a complex distribution of  
87 sulfur and oxidized sulfur compounds. Based on a similar approach, Lu et al. (2014) detected  
88 high levels of alkylcyclothioethers within Jiangnan Basin oils. More recently, Liu et al. (2018)  
89 reported an in-depth investigation of polar sulfur compounds in immature crude oils from the  
90 Jiangnan Basin via ESI-N and ESI-P FTICR-MS and showed that the detected S-bearing species  
91 are mostly cyclic, originating from intramolecular sulfurization of functionalized precursors  
92 during early diagenesis. The authors inferred that the extent of intramolecular sulfurization is  
93 based on the number of reactive functional groups in the precursor molecule.

94 Expanding on these findings, the present study investigates the compounds present in a suite of  
95 sulfur rich oils and their fractions via APPI-P FTICR-MS and APPI-P FTICR-MS/MS  
96 experiments. Extended compositional information of high molecular weight, sulfur bearing  
97 molecules that could not be detected in earlier GC based studies is presented and critically  
98 discussed based on the current understanding of sulfur incorporation into organic matter. This  
99 paper aims to gather new insights from sulfur rich oil molecular composition to probe the  
100 mechanisms of sulfur incorporation into sedimentary organic matter, which could possibly be  
101 leveraged for the development of carbon dioxide removal (CDR) technologies.

## 102 **2. Materials and methods**

### 103 *2.1. Sample set description*

104 Five sulfur rich oils from different basins were selected in this study: (a) three high maturity,  
105 highly biodegraded oils from the Peace River area, Canada; (b) one immature oil from the  
106 Jiangnan Basin, Eastern China; and (c) one immature oil from the Rozel Point, Utah, USA  
107 (Table 1).

108 Rozel Point is an immature, sulfur-rich heavy oil generated from a hypersaline, lacustrine source  
109 rock of Miocene age (Meissner et al., 1984; ten Haven et al., 1988; Sinninghe Damsté et al.,  
110 1989). High concentrations of organic sulfur (up to 15 wt%) and the unique depositional  
111 environment that characterizes Rozel Point oil have motivated many studies of the origin and  
112 fate of organically bonded sulfur during early diagenesis, with implications for the understanding  
113 of Type II-S kerogen formation (Eglinton et al., 1994). The Jiangnan Basin, located in eastern  
114 China, is comprised of five major tectonic units with the Qianjiang depression being the most  
115 significant structure since most oil production takes place in this region from the Eocene  
116 Qianjiang Formation (Philp and Zhaoan, 1987). Oils from the Qianjiang depression are sourced  
117 from the Qianjiang Formation and Xingouzui Formation, which were deposited in anoxic, sulfate  
118 reducing and saline lacustrine environments during the early Cretaceous to Paleogene (Carroll  
119 and Bohacs, 2001; Hou et al., 2017). Due to cycles of marine transgression and progradation  
120 activity, hot paleoclimate, and clastic sediment deposition from lacustrine systems, over 220  
121 evaporitic layers interlaced with shales and sandstones formed the Jiangnan Basin source rocks  
122 (Philp and Zhaoan, 1987; Peirong et al., 2008; Hou et al., 2017). The Qianjiang Formation is  
123 suspected to act as both the source rock and reservoir (Philp and Zhaoan, 1987; Huang and

124 Hinnov, 2014; Hou et al., 2017). Upper Qianjiang sections act as the oil reservoir while the  
125 deeper sections and Xingouzui Formation embody the source rock. Oils from the Qianjiang  
126 Formation typically exhibit a pronounced even/odd *n*-alkane predominance, low ratios of Pr/Ph  
127 and high ratios of gammacerane/hopane, indicating a highly anoxic, reducing and saline  
128 depositional environment (Philp and Zhaoan, 1987). The Peace River oil sands, located in  
129 northwestern Alberta, is one of three major bitumen deposits in that province. The Peace River  
130 area is comprised of two main groups, the Bullhead Group formed by the Cadomin and Gething  
131 formations, and the overlying Fort St. John Group which, in succession, comprises the Bluesky,  
132 Spirit River, Peace River and Shaftesbury formations. Peace River oil reservoirs, Bluesky,  
133 Gething and McMurray, are late Paleozoic to Mesozoic in age and are charged by Cretaceous oil  
134 source rocks (Adams et al., 2012). The reservoirs are supplied with a mixture of hydrocarbons  
135 expelled from multiple source rocks including the Nordegg Member of Fernie Formation,  
136 Gordondale Formation and Permian Doig Formation (Riediger, 1994; Adams et al., 2013).  
137 Manville Group reservoirs, Bluesky and Gething, may also be vertically charged by the  
138 Mississippian Exshaw-Banff formations where hydrocarbons migrated upwards to permeable  
139 Cretaceous sand units when erosion deteriorated the seal enforced by the Poker Chip shale (Allan  
140 and Creaney, 1991). Peace River bitumen, which exhibits high levels of biodegradation, such as  
141 those from the Bluesky Formation, are sourced by Exshaw - Banff source rock as these oils had a  
142 longer residence time in the reservoir. On the other hand, deposits in the western part of Peace  
143 River received more significant contributions from Gordondale source rock as these oils are  
144 higher in sulfur content, less mature and have greater API gravity (Adams et al., 2013).

145

## 146 *2.2. Non-polar macromolecular fractions*

147 The experimental procedure used to obtain the non-polar macromolecular fraction (NPMF) from  
148 Rozel Point and Jiangnan oils was adapted from Adam et al. (1993). In brief, aliquots of the  
149 crude oils (~100 mg) were loaded on a silica gel (6.0 g, 220-440 mesh) column. The samples  
150 were washed with 50 mL of hexane and then eluted with 15 mL of hexane:diethyl ether (98:2),  
151 which yielded the NPMF orange-colored band that was collected, evaporated to dryness and  
152 submitted to APPI-P FTICR-MS analysis. NPFM was not produced from western Canadian oils  
153 indicating that the macromolecular fraction was not present in such mature and overmature oils.

154        *2.3. Mass spectrometry*

155        APPI-P is the ionization technique of choice for sulfur-rich sample analysis due to its enhanced  
156        ability to ionize sulfur compounds in complex organic mixtures without chemical derivatization  
157        nor special sample preparation strategies (Purcell et al., 2006; Oldenburg et al., 2014). Two  
158        primary ionization mechanisms are expected to occur in the APPI-P source: proton transfer,  
159        which generates even electron protonated ions depending on the proton affinity of the analyte;  
160        and the formation of an odd electron radical ion species (Raffaelli and Saba, 2003). Note that the  
161        thermospray conditions in APPI-P show negligible in-source thermal alteration of species  
162        (Bagag et al., 2008).

163        Even though it was demonstrated in previous studies that APPI-P FTICR-MS has some very  
164        rudimentary quantitation capabilities (Oldenburg et al., 2014), due to unassessed differences in  
165        ionization responses, intensities cannot be assumed to reflect the actual abundance of compounds  
166        present in the sample. Thus, quantitative aspects of compounds and compound classes are not  
167        considered herein; only relative intensities are discussed, based on monoisotopic peak intensities  
168        of assigned peaks (RMI, relative monoisotopic intensity). Also, structural discussions of detected  
169        species are only speculative, since FTICR-MS data provide unambiguous molecular formulas but  
170        cannot distinguish isomers.

171        Attempts to perform FTICR-MS/MS experiments in complex mixtures provide limited but  
172        potentially useful information. Targeted peaks may represent a dozen to potentially thousands of  
173        different isomers, as no prior chromatographic clean up was performed in these experiments.  
174        Therefore, the measured fragmentation patterns represent the conflation of the fragmentation  
175        patterns of all the isomers with the targeted molecular formula. The isolation window ( $m/z \pm 0.4$ )  
176        used in this study may contain 10+ peaks. Thus, only peaks with the highest intensity within  
177        their  $m/z \pm 0.4$  spectral windows were selected for MS/MS experiments. To avoid interferences  
178        with fragments originating from other parent peaks within the  $m/z \pm 0.4$  windows, detected peaks  
179        with more carbon or sulfur atoms than the parent ion were ignored. Peaks detected in MS/MS  
180        mode were assigned molecular formulae with assignment errors < 100 ppb.

181        *2.3.1. APPI-P FTICR-MS method*

182 The samples were analyzed using a 12 T Bruker Solarix FTICR-MS instrument. Whole oil and  
183 fractions were diluted to 0.25 mg/mL in toluene and then infused into the ionization source using  
184 a syringe pump set to deliver 200  $\mu\text{L}/\text{h}$ . Atmospheric pressure photoionization in positive ion  
185 mode via a krypton lamp at 10.6 eV was used to ionize the samples. Transfer capillary  
186 temperature and nebulizer pressure were set to 350  $^{\circ}\text{C}$  and 1.0 bar, respectively. Reserpine  
187 ( $\text{C}_{33}\text{H}_{40}\text{N}_2\text{O}_9$ ) was added to the sample solution to assess internal calibration efficiency. The  
188 instrument was tuned using a reference Athabasca whole bitumen sample. Ions ranging from  $m/z$   
189 150 to 1500 were isolated using a linear quadrupole and accumulated over 5 ms in the collision  
190 cell before being transferred to the ICR cell. Spectra were collected in absorption mode, using an  
191 algorithm proposed elsewhere (Kilgour et al., 2013). Two hundred transients of 8 million points  
192 in the time domain were collected and summed to improve the experimental signal/noise ratio  
193 (SNR). FTICR-MS raw data were processed using the CaPA v.1.0 (Aphorist Inc.) software  
194 package. Peaks with SNR higher than 4 were assigned based on highly accurate  $m/z$   
195 measurements and on stable isotopic distributions. Compositional boundaries, in terms of  
196 stoichiometry (element atom content) for the fitting algorithm, were set to  $\text{C}_{4-95}\text{H}_{0-200}\text{N}_{0-2}\text{S}_{0-8}\text{O}_{0-}$   
197  $_5$  and double bond equivalent (DBE) range, which is a measure of hydrogen deficiency due to  
198 double bonds and/or cyclic structures, was limited to be between 0 and 60. The mass spectra  
199 were recalibrated using homologous series present in the samples. Ragnarök v.2.0 (Aphorist Inc.)  
200 was used for data manipulation and visualization.

### 201 2.3.2. APPI-P FTICR-MS/MS method

202 A set of peaks found in the oils were selected for MS/MS experiments (Sections 4.3, 4.4 and  
203 4.7). Sulfur-bearing parent ions were isolated ( $m/z \pm 0.4$ ) in the linear quadrupole and  
204 accumulated over 0.3–1.0 s in the collision cell. The collision cell voltage was set to 12–20 V  
205 (compared to regular operating levels at 5 V) to achieve collision-induced dissociation (CID)  
206 conditions, causing the ions to undergo several low energy collisions with room temperature  
207 argon atoms. Ions in the collision cell (parent and daughters) were then transferred to the ICR  
208 cell and the MS/MS spectra were collected in absorption mode after summing two hundred  
209 transients of 4 million points in the time domain. The instrument was tuned to detect ions  
210 ranging from  $m/z$  150 to 1500, i.e. small molecular fragments ( $< m/z$  150) fell outside the



211 analytical window in these experiments. Bruker DataAnalysis v.4.4 and R v.3.4.0 software  
212 packages were used for data processing and visualization.

### 213 **3. Results**

#### 214 *3.1. NPMF fractionation*

215 The NPMF fractions represented 18 and 22 wt % of the Jiangnan and Rozel Point oils,  
216 respectively, and their APPI-P mass spectra are shown in Figure 1.

#### 217 *3.2. APPI-P FTICR-MS of oils and their fractions*

218 In all spectra acquired in this study, the isotopic pattern of  $^{13}\text{C}$  peaks indicated the presence of  
219 singly charged species only. Averaged mass resolution ( $m/\Delta m 50\%$ ) achieved for peaks detected  
220 between  $m/z$  397 and 403 was higher than 1,015,000 and higher than 560,000 for peaks between  
221  $m/z$  797 and 803. Unique molecular formulas with absolute assignment error lower than 750 ppb  
222 were assigned to 40,967 monoisotopic peaks present in the oil samples spectra (Fig. 1). In each  
223 sample, peaks left unassigned represented less than 3% of total peak intensity and are not  
224 discussed further. Heteroatom classes with fewer than 25 peaks were also excluded from the  
225 discussion.

226 Fig. 2 shows the heteroatom class distribution of the investigated oils, as obtained by APPI-P  
227 FTICR-MS, where radical ion classes are differentiated from protonated ones by a dot following  
228 the class symbol. The relative sulfur number distribution for the studied oils, as measured in  
229 classes  $\text{S}\bullet_{1-6}$ , is shown in Fig. 3A. Figs. 3B and 3C show the carbon number and DBE  
230 distribution of all detected molecular formulae within the oil set. Fig. 4 highlights the DBE  
231 distribution within the sulfur containing heteroatom classes, whereas Fig. 5 shows the carbon  
232 number distribution of heteroatomic classes  $\text{S}\bullet_1$  DBE 5 and 6 species. Fig. 6 shows the carbon  
233 number distribution of heteroatomic classes  $\text{S}_1$  DBE 1 and  $\text{S}\bullet_1$  DBE 3. Fig. 7 shows the DBE  
234 distribution of  $\text{C}_{40}$  species within class  $\text{S}\bullet_{1-5}$  as found in Rozel Point and Jiangnan oils, whereas  
235 Fig. 8 shows the heteroatomic class  $\text{S}\bullet_1$  DBE 8 carbon distribution for the same oils.

236

#### 237 *3.3. APPI-P FTICR-MS/MS*

238 CID experiments were set to investigate the fragmentation patterns of a few selected parent ions  
239 by measuring fragments with  $m/z > 150$ . Fig. 9 shows a representation of the APPI-P FTICR-  
240 MS/MS spectra of class  $S_1$  DBE 5  $C_{27-30}$  species analyzed in the JH and RP oils, whereas Fig. 10  
241 highlights data obtained from APPI-P FTICR-MS/MS spectra of class  $S_{\bullet 2}$  DBE 11  $C_{56-57}^{\#}$   
242 species.

## 243 **4. Discussion**

### 244 *4.1. NPMF and sulfur-bridged (macro)molecules*

245 The current understanding of organic matter sulfurization mechanisms assumes the substitution  
246 of functional groups in the precursor molecules by sulfur atoms, which then can rearrange to  
247 more stable forms such as thiolanes and thiophenes, and/or act as bridges linking two precursor  
248 molecules. As the diagenesis progresses, such linkages would propagate and eventually generate  
249 a Type II-S kerogen, which can produce oil molecules at a much lower level of thermal stress  
250 than other kerogen types (Tegelaar et al., 1989; Vandenbroucke and Largeau, 2007).

251 Adam et al. (1993) reported a hexane soluble nonpolar macromolecular fraction (NPMF)  
252 occurring in sulfur rich oils (including the RP oil), composed of highly aliphatic, high molecular  
253 weight structures, possibly cross-linked by sulfur in a process similar to natural vulcanization.  
254 Through a stepwise selective sulfur removal procedure, Richnow et al. (1993) proposed that in  
255 RP oil, *n*-alkanes, hopanoids, steroids and phytanes are bound simultaneously via oxygen, sulfur  
256 and aromatic units, based on the position of the functionality in the precursor molecules, creating  
257 a macromolecular matrix. Alcohols released in the desulfurization of RP oil NPMF can be as  
258 abundant as hydrocarbons; they exhibited similar carbon number distribution to hydrocarbons,  
259 but the origin of most species could not be assigned (Jenisch-Anton et al., 1999). Efforts to  
260 chemically or physically degrade such macromolecular organic matter were needed to  
261 circumvent the inefficiency of gas chromatography in analyzing sulfur rich macromolecules  
262 (Jenisch-Anton et al., 1999). Investigating the RP oil, Adam et al. (1993) reported the presence  
263 of a NPMF composed of cross-linked molecules containing multiple sulfur atoms, which  
264 represented up to 32 wt% of the whole material (Adam et al., 1992; Jenisch-Anton et al., 1999).  
265 Significantly different average molecular weight values of NPMF from RP oil were reported  
266 using various techniques (Adam et al., 1993): 815 Da by vapor pressure osmometry; 3400 Da by  
267 light scattering measurements and 1660 Da by size exclusion chromatography.

268 Despite a minor clean up of non-S-bearing heteroatom classes that was applied during sample  
269 preparation, FTICR-MS spectra of NPMF fractions of JH and RP oils analyzed in this study are  
270 very similar to the respective whole oil spectra (compare panels A vs. B and C vs. D in Fig. 1),  
271 indicating that the analyses of whole oils by APPI-P are representative and capture the major S-  
272 bearing constituents in those oils.

273 The spectra shown in Fig. 1 (A-G) revealed no detectable peaks above  $m/z$  1100 in all the  
274 analyzed oils and fractions. To further corroborate the absence of high molecular weight  
275 compounds, we focused on RP oil, as illustrative of very low maturity oils, and because it has  
276 been well studied previously. An APPI-P FTICR mass spectrum was collected from  $m/z$  1000–  
277 3000 (Fig. 1H), after tuning the instrument with Agilent Low Concentration Tuning mix for  
278 enhanced sensitivity at this  $m/z$  range. A group of peaks was detected at  $m/z$  1000-1250 in Fig.  
279 1H. Yet, no evidence could be found of larger molecules ( $m/z > 1500$ ) in the RP oil. Considering  
280 that the NPMF is putatively formed by consecutively adding building block units, in a process  
281 similar to vulcanization, the detection of fragments at  $m/z > 1250$  would be expected since a  
282 molecular size continuum should exist (Eglinton et al., 1994). The Jiangnan and Rozel Point oil  
283 NPMF spectra (Fig 1B and 1D, respectively) also showed no indication of ‘macromolecules’ in  
284 their composition. Noteworthy, there is no evidence of in-source fragmentation, consistent with  
285 previous work by Bagag et al. (2008) on the APPI analysis of sensitive biomolecules, nor  
286 evidence of fragmentation during the ion transfer to the ICR cell. In addition, the samples fully  
287 dissolved in toluene and no evidence of molecular aggregation was seen visually or in the mass  
288 spectra. Hence, the APPI-P FTICR-MS spectra obtained herein represent a key piece in the  
289 understanding of the so called NPMF. Clearly, the APPI-P FTICR-MS results indicate that there  
290 might not be a macromolecular fraction present in RP and JH oils at all.

291 Intermolecular sulfur incorporation is inferred from the increased RMI detected in the  $C_{54-59}$   
292 range within class  $S_{\bullet 1}$  DBE 8 (Fig. 8), tentatively interpreted to be the indication of two  $C_{27-29}$   
293 steroidal units bridged by one sulfur atom. Noteworthy, a similar pattern was seen in class  $S_{\bullet 2-3}$   
294 DBE 9-11  $C_{54-59}$  species, but not in class  $S_{\bullet 1}$  DBE 7. RP class  $S_{2\bullet}$  DBE 11  $C_{56-57}$  compounds  
295 were selected as targets for MS/MS experiments (Fig. 10) and results show detectable daughter  
296 ions with the loss  $C_{28}$  and  $C_{29}$  with no sulfur and  $\Delta$ DBE of -5.5, suggestive of an intermolecular  
297 linkage.

298 Although inferred, intermolecular sulfur incorporation reactions that yield high molecular weight  
299 molecules are limited. Recent investigations have recognized the impact of nanoaggregation on  
300 molecular weight measurements of other oil components, such as asphaltenes (Zhang et al.,  
301 2013). Initially thought to be as high as several thousand kilodaltons, the average molecular  
302 weights of non-aggregated asphaltene distributions are now recognized to be around 750 Da  
303 (Mullins, 2010; Hosseini-Dastgerdi et al., 2015; Snowdon et al., 2016). Similarly, the  
304 aggregation of sulfurized species might be the reason for the overestimation of NPMF molecular  
305 weights in previous studies.

#### 306 *4.2. Thermally immature vs. biodegraded sulfur rich oils*

307 Sinninghe Damsté and De Leeuw (1987) used gas chromatography coupled to mass spectrometry  
308 (GC-MS), to identify not only sulfur containing isoprenoids (C<sub>15</sub> and C<sub>20</sub>) in Rozel Point oils but  
309 also a series of isoprenoid chains bonded to several organosulfur structures (e.g. thiophenes,  
310 thiolanes, benzothiophenes). Further work led by the same authors expanded the scope of sulfur  
311 compounds detected by GC-MS: C<sub>30</sub> and C<sub>35</sub> isoprenoid thiophenes, alkylthianes, isoprenoid  
312 thiolanes, thiophene steranes, thiolane steranes, alkylbenzothiophenes, isoprenoid  
313 benzothiophenes, isoprenoid bithiophenes and sulfur containing hopanoids. In total, around one  
314 thousand sulfur compounds were identified, with a molecular weight upper limit of 600 Da and a  
315 maximum of two sulfur atoms in the structure (Sinninghe Damsté et al., 1987). Such species  
316 were considered different from those sulfur compounds found in most oils since their structures  
317 are closely related to biogeochemical precursors (Sinninghe Damsté et al., 1989). Lu et al. (2014)  
318 investigated a set of Jiangnan oils using ESI-P ion mode, after S-methylation reactions and  
319 detected high levels of alkyl cyclothioethers within the oils. The identified peaks ranged from  
320 classes S<sub>1-3</sub>, DBE 1-14, and C<sub>10-35</sub>. Compounds from class S<sub>1</sub> DBE 1-3 were detected with  
321 remarkably higher relative intensities. The detected predominance of C<sub>20-21</sub> compounds was  
322 attributed to the sulfurization of phytanic acids and phytols, and the overall sulfurization  
323 mechanism in Jiangnan oils was hypothesized to involve carboxylic acids and fatty alcohols,  
324 which resulted in the formation of cyclic thioethers. Kohnen et al. (1993) identified sulfur-bound  
325 steroid and phytane moieties in the Jiangnan oil (also in Rozel Point oil), suggesting that di- or  
326 polysulfide linkages are present in S-containing moieties and the position of double bonds in the  
327 precursors controls the position of S-linkages. Liu et al. (2018) showed that the ubiquitous sulfur

328 ring structures present in the organic sulfur compounds in Jiangnan Basin oils likely originated  
329 from intramolecular sulfurization reactions.

330 The biogeochemistry of Peace River oils has been extensively studied in our group by Adams et  
331 al. (2013) and Bennett et al. (2013). More recently, Oldenburg et al. (2017) have shown the  
332 effects of different biodegradation levels in the chemical composition distribution in a Peace  
333 River Area, Bluesky Formation reservoir profile, as measured by FTICR-MS. Overall, molecules  
334 with higher DBE and higher sulfur number were found to be more resistant to biodegradation.

335 The results of this study show that the mass spectra of Jiangnan and Rozel Point oils are  
336 remarkably different from the spectra of the Peace River oils (Fig. 1, A and C vs. E – G). The  
337 smoother distribution of peaks detected in the Peace River oils, similar to a bell-shaped curve, is  
338 a very common observation in fossil fuel analysis (Marshall and Rodgers, 2008) and relates to  
339 the loss of carbon number preferences caused by the thermal cracking of kerogen during oil  
340 generation. On the other hand, both Jiangnan and Rozel Point oils show mass spectra that reflect  
341 a dominant signature of precursor compounds (Fig. 3C), somewhat akin to the spectra observed  
342 in recent marine sediments (Radović et al., 2016a), a consequence of the low thermal stress  
343 levels experienced by these oils. Thus, both Rozel Point and Jiangnan oil APPI-P FTICR-MS  
344 mass spectra are a good representation of the chemical species and mechanisms involved in early  
345 diagenesis and related sulfur incorporation processes.

346 In these two oils, the predominance of heteroatom classes with one or more sulfur atoms ( $S_{1-6}$ ) is  
347 evident (Fig. 2), whereas classes HC,  $O_{1-2}$ , and  $NO_{0-1}$  are the only non-sulfur containing  
348 heteroatom classes detected, accounting in total for less than 10% RMI in JH and RP oils. This is  
349 the first report of species with more than two sulfur atoms per molecule in Rozel Point oils,  
350 while Liu et al. (2018) detected up to  $S_5$  compounds in Jiangnan oils.

351 The Peace River oils exhibit higher RMI for  $S_{1\bullet}$  species, and a decreased RMI as the sulfur  
352 number increases up to  $S_{4\bullet}$  (Fig. 3A). Both Rozel Point and Jiangnan oils, on the other hand,  
353 show maximum RMI at  $S_{2\bullet}$  and extend up to  $S_{6\bullet}$  and  $S_{5\bullet}$ , respectively. Such differences may be  
354 associated with the depositional settings favoring organic matter reaction with (poly)sulfide,  
355 coupled to the low thermal stress experienced by the Jiangnan and Rozel Point oils. Therefore,  
356 the relative ratios of lesser and more extensively sulfurized compound classes (e.g.  $S_{1\bullet}$  vs.  $S_{2\bullet}$ ) is

357 proposed as a potentially robust proxy for source rock depositional settings, although further  
358 testing with an extended sample set is required.

359 Peace River oils WC and BS show very similar chemical composition (Figs. 2 and 3). Despite  
360 showing a similar carbon number distribution to its Peace River analogs in Fig. 3C, the Peace  
361 River GR oil DBE distribution is significantly shifted towards lower DBE values (Fig. 3B),  
362 while exhibiting a relative enrichment in  $S_2\bullet$  species (Fig. 3A). Geochemical differences among  
363 the Peace River oils studied herein have been observed before by Adams et al. (2013), including  
364 a lower thermal maturity, lower biodegradation extent and higher sulfur content for oils primarily  
365 sourced from the sulfur enriched Gordondale Formation.

#### 366 *4.3. Sulfurized steroids and hopanoids (class $S_1\bullet$ DBE 5-6)*

367 In a study by Lu et al. (2013),  $C_{28-30}$  DBE 5-7 steroids were detected in class  $S_1$  of a sulfur rich  
368 heavy oil in Jiaxian Sag, Bohai Bay Basin, China. Sulfurized steroidal structures were previously  
369 reported in RP oil as part of the macromolecular matrix through a sulfur linkage located in ring  
370 A or B (Adam et al., 1992). Thiophene and thiolane steranes have been identified in RP oil by  
371 Sinninghe Damsté et al. (1987). Using selective cleavage of acyclic sulfide by superheated  
372 methyl iodide, Schouten et al. (1993) described the release of  $C_{27-30}$  steroid products from RP  
373 polar and asphaltene fractions. Kohnen et al. (1993) suggested that multiple substrates prone to  
374 sulfur incorporation reactions, such as  $\Delta^2$ -,  $\Delta^3$ -,  $\Delta^5$ -sterenes or  $\Delta^3,5$ -steradienes, give rise to a  
375 large variation in the positions and stereochemistry of sulfur incorporated steroids and that  
376 timing of sulfur incorporation during diagenesis would significantly influence the resulting  
377 products.

378 As can be observed in both in the overall (Fig. 3B) and class  $S_1\bullet$  (Fig. 4A) DBE number  
379 distribution plots, there is a high RMI of DBE 5-6 species in Rozel Point and Jiangnan oils.  $C_{27-}$   
380  $_{30}$  species largely dominate the class  $S_1\bullet$  DBE 5 carbon number distribution (Fig. 5A). Since  
381 DBE 4 species do not show an evident RMI increase for  $C_{27-30}$  in any compound classes (except  
382 for  $HC\bullet$ ), sulfurized steroid-like compounds likely include a cyclothioether moiety to result in a  
383 DBE 5 value (Fig. 5A). Although present,  $C_{27-30}$  species do not exhibit an increased RMI within  
384 heteroatom classes with more than 2 sulfur atoms. This suggests that less functionalized  
385 molecular precursors such as steroids yield less sulfur altered diagenesis products, in contrast to

386 precursors such as carotenoids that have multiple reactive sites for sulfur incorporation (Section  
387 4.5).

388 The prominent class  $S_1 \bullet$  DBE 6  $C_{35}$  peak (Fig. 5B) may represent S-bound  $C_{35}$  homohopanes,  
389 which are known to be the dominant triterpene in the Rozel Point oil. The increased RMI of class  
390  $S_1 \bullet$  DBE 6  $C_{27-29}$  species in RP oil is remarkable (Fig. 5B) because there have been no reports of  
391 S-bound norhopanes despite the occurrence of S-bound (homo)hopanes in thermally immature  
392 sedimentary organic matter. Within the most plausible routes for the production of sulfurized  
393 steroids in natural environments (Lu et al., 2013), none would yield S-steroid species with DBE  
394 6. To further investigate the forgoing interpretations, ions from class  $S_1 \bullet$  DBE 5  $C_{27-30}$  were  
395 selected as targets for MS/MS experiments (Fig. 9). In both RP and JH oils, detected daughter  
396 ions represent the loss of  $C_{1-6}$  fragments with  $\Delta$ DBE -0.5, i.e. aliphatic moieties such as  $-CH_3$ , -  
397  $C_2H_5$ , and no sulfur atoms, compatible with a side carbon chain fragmentation from a D-ring or  
398 an AB-ring sulfurized steroid (Lu et al., 2013). The RP oil targeted ions show, although with low  
399 relative intensity,  $\Delta C_{0-2} \Delta S_1$  daughter ions, which indicate that some class  $S_1 \bullet$  DBE 5  $C_{27-30}$   
400 isomers might have thiol ( $\Delta C_0 \Delta S_1$  fragments) or aliphatic sulfide ( $\Delta C_{1-2} \Delta S_1$ ) as functional  
401 groups, instead of cyclic sulfide. In such cases, the parent DBE 5 ion could represent, among  
402 other structures, either a 5-ring moiety (e.g. hopanoids) or a 4-ring structure with one carbon-  
403 carbon double bond (e.g. sterenes). However, to the best of our knowledge, no sterenes have  
404 been detected in RP oil, despite being thermally immature.

#### 405 *4.4. Class $S_1$ DBE 1 and class $S_1 \bullet$ DBE 3 species*

406 Previous work from Lu et al. (2014) indicated that the Jiangnan oils are relatively enriched in  
407 Class  $S_1$  DBE 1-3 species, and the same pattern was observed herein (Figs. 4A,B). The high  
408 RMI of class  $S_1$  DBE 1 species suggests an enrichment in thiolanes or thianes, whereas high  
409 RMI of class  $S_1 \bullet$  DBE 3 species suggest thiophenic structures. There is an even/odd  
410 predominance spanning  $C_{16-40}$  class  $S_1$  DBE 1 species in Jiangnan oil (Fig. 6A). Lu et al. (2014)  
411 reported the occurrence of odd/even predominance in the Jiangnan oil class  $S_1$ , centered around  
412 the  $C_{21}$  peak. Although not discussed in their paper, the methylation reactions used by Lu et al.  
413 (2014) to facilitate ESI-P detection of sulfur species added one carbon to all the species detected  
414 in their study, and this might have caused the reversal of actual even/odd carbon preference to  
415 observed odd/even preference in their data. In agreement with our observations, Sinninghe

416 Damsté et al. (1987) and Sheng et al. (1987) reported a series of C<sub>10-32</sub> alkylthiolanes and  
417 alkylthianes in Rozel Point and Jiangnan oils, respectively, which also exhibited a strong  
418 even/odd preference. Herein, the APPI-P results show that the even/odd preference extends much  
419 further, up to C<sub>40</sub> in the Jiangnan oil (Fig. 5A).

420 CID experiments were set to investigate the fragmentation patterns of class S<sub>1</sub> DBE 1 C<sub>26,28,30</sub>  
421 species. However, the experiments failed to produce any detectable daughter ion at  $m/z > 150$ .  
422 Typically, CID promotes the rupture of the weakest bonds in a molecule, therefore the  
423 intermediate species generated after C-S bond cleavage in thiolanes/thianes might have  
424 undergone extensive fragmentation. Based on the same reasoning, the presence of aliphatic thiols  
425 or dialkyl sulfides contributing to class S<sub>1</sub> DBE 1 species cannot be precluded, although the lack  
426 of class S<sub>1</sub> DBE 0 species indicates such functional groups may be largely absent in the oil  
427 matrices. Initially thought as intermediates in the formation of thiophenes, thiolanes have been  
428 detected in sediment extract but their association with thiophene formation is yet unclear  
429 (Sinninghe Damsté et al., 1986). The APPI-P FTICR-MS results support the hypothesis that (a)  
430 these cyclic sulfides, formed during very early stages of diagenesis, likely never partake in the  
431 kerogen formation (Brassell et al., 1988; Peng et al., 1998); and that (b) these cyclic sulfides may  
432 be derived from the same functionalized precursors as the *n*-alkanes, such as *n*-alkanoic acids  
433 and *n*-alkanols (Brassell et al., 1988; Liu et al., 2018).

434 Class S<sub>1</sub>• DBE 3 species, which putatively represent thiophenic structures, show increased RMI  
435 at C<sub>20,30,35,40,45</sub> in the Rozel Point oil, and at C<sub>20,24-26,28,30,40</sub> in the Jiangnan oil (Fig. 5B). FTICR-  
436 MS data alone is not capable of distinguishing isomers, but the increased RMI of multiple  
437 isoprene units (x•C<sub>5</sub> units) may indicate the detection of an extended range of isoprenoid  
438 thiophenes. To support this reasoning, isoprenoid thiophenes with < 35 carbons have been  
439 identified by GC-MS in Rozel Point and Jiangnan oils, as reported by Sinninghe Damsté et al.  
440 (1987) and Sheng et al. (1987), respectively.

#### 441 4.5. C<sub>40</sub> species

442 The overall carbon distribution plot (Fig. 3C) shows a high RMI for C<sub>40</sub> species in both RP and  
443 JH oils. This observation alone suggests the presence of compounds of biological provenance  
444 rather than produced by thermal cracking of kerogen, which typically smooths out any carbon



445 number preference (e.g. even/odd preference). Carotenoids are the likely biological source for  
446  $C_{40}$  species, since they are common constituents of several living organisms, from archaea and  
447 cyanobacteria to higher plants and animals (Walter and Strack, 2011). Recently, we have used a  
448 similar APPI-P FTICR-MS approach to characterize carotenoids in recent marine sediments, in  
449 addition to various other lipid markers (Radović et al., 2016a). Whereas  $C_{27-30}$  S-bearing steroids  
450 appear to be limited to classes  $S_{1-2}$  (see Section 4.3.),  $C_{40}$  S-carotenoids are seen in classes with  
451 up to six sulfur atoms. Such findings indicate the clear dependence of the degree of sulfur  
452 incorporation on the number of double bonds present in the lipid undergoing diagenesis, and  
453 support the observations and interpretations reported previously with ESI FTICR-MS data (Liu  
454 et al., 2018). Carotenoids can offer multiple spots for sulfur incorporation, contrary to steroids  
455 which have a limited number of double bonds and oxygenated groups. The predominance of  $C_{40}$   
456 in classes  $S_{5-6}$  is remarkable and reveals that, overall, most of the sulfurized species may  
457 resemble the biological precursors in both Jiangnan and Rozel Point oils.

458 Typical carotenoid structures (e.g., lycopene,  $\beta$ -carotene, nostoxanthin) have 13 DBE units,  
459 although higher DBE species are also possible (Walter and Strack, 2011). Carotenoid diagenesis  
460 in different depositional settings can be quite complex (Watts and Maxwell, 1977; Repeta and  
461 Gagosian, 1987), but the identification of several reduced carotenoids in recent sediments  
462 suggests that hydrogenation occurs during diagenesis without structural or stereochemical  
463 specificity (Hebting et al., 2006). Adam et al. (1993) not only suggested acyclic carotenes,  $\beta$ -  
464 carotene and monocyclic carotenes as building blocks of Rozel Point NPMF, but also discussed  
465 their implications to organic matter input determination. In the current study, similarly to Liu et  
466 al. (2018), sulfurized carotenoids are present in the oils as free molecules instead of being  
467 building blocks of NPMF, as evident by the absence of cross-linked carotenoid-derived NPMF  
468 intermediates (i.e., no significant signal from  $C_{80}$  species was detected). In summary, the sulfur  
469 incorporation in carotenoid precursors is not followed by an increase in the carbon number, thus  
470 mostly original (biological)  $C_{40}$  species with varying DBE values were detected.

471 Fig. 8 shows the DBE distribution of classes  $S_{1-5}$   $C_{40}$  species. A significant pattern observed in  
472 Fig. 8 is the increased RMI for the  $DBE = S + (2,4,6)$  species, which might reflect the level of  
473 ‘thiophenization’. That is,  $DBE = S + 0$  species would represent altered aliphatic carotenoids  
474 (e.g. lycopene) where the sulfur atoms are incorporated as thiolane.  $DBE = S + 2$  species

475 represent the altered carotenoids with 2 cyclic structures (e.g. beta-carotane) where sulfur atoms  
476 are incorporated as thiolane. Note the higher RMI of DBE S + 2 compared to DBE S + 0 species,  
477 reflecting the higher abundance of cyclic carotenoid structure in the deposited organic matter.  
478 Since the conversion of a thiolane into thiophenes involves a  $\Delta$ DBE +2 shift, species with DBE  
479 S + 4 and S + 6 may reflect the number of thiophenic structures present in the compound. The  
480 transformation of thiolanes into thiophenes is likely dependent on the thermal stress level  
481 experienced by the sample. The ratios of thiophene and thiolane structures within the carotenoid  
482 classes require further investigation to determine its potential usefulness as a marker for early  
483 diagenesis sulfurization.

#### 484 *4.6. The DBE number $\geq S_x$ rule*

485 In Fig. 8, the minimum detected DBE number in each plot is equal to the number of sulfur atoms  
486 of the corresponding class, i.e. DBE number  $\geq S_x$ , which indicates that, for C<sub>40</sub> assignments, the  
487 sulfur incorporation reactions are associated with a DBE increase. In fact, such a pattern is  
488 observed in the dominant molecular formulae detected in this study, indicating that polysulfides,  
489 aliphatic sulfides and thiols must be largely absent as stand-alone functional groups in the RP  
490 and JH oils. The only exceptions to the rule are class S<sub>2</sub>• DBE 1 C<sub>16,18,20,22,24,26</sub> and class S<sub>3</sub>•  
491 DBE 2 C<sub>37</sub> species. The even/odd predominance in class S<sub>1</sub> DBE 1 was discussed previously.  
492 The species in class S<sub>2</sub>• DBE 1 C<sub>16,18,20,22,24,26</sub> probably relates to those where an additional  
493 sulfur atom is incorporated as a cyclic polysulfide or a thiol. The class S<sub>3</sub>• DBE 2 C<sub>37</sub> species are  
494 discussed in Section 4.7.

#### 495 *4.7. Sulfurized alkenones*

496 Alkenones are long chain unsaturated ketones produced by some phytoplankton species and  
497 typically used as paleoenvironmental proxies (Volkman et al., 1980; Brassell et al., 1986). An  
498 increased RMI was observed for classes S<sub>1-3</sub>• DBE 1-3 C<sub>37</sub> species in Rozel Point oil, suggesting  
499 that alkenones have also undergone sulfurization during diagenesis. Although it is tempting to try  
500 and estimate the ratios of parent C<sub>37:2</sub> and C<sub>37:3</sub> (e.g. UK`37) based on their sulfurized analogs,  
501 this exercise would be highly speculative at this point. Also, the occurrence of class S<sub>3</sub>• DBE 2  
502 C<sub>37</sub> species as an exception to the DBE number  $\geq S_x$  rule, indicates that in some cases thiol or  
503 polysulfides might be present. The peak representing class S<sub>2</sub>• DBE 2 C<sub>37</sub> in Rozel Point oil was

504 selected for MS/MS experiments, and similarly to class S<sub>1</sub> DBE 1 peaks, no daughter ions could  
505 be detected, except for an ion representing the loss of a -SH group. This result suggests that the  
506 C<sub>37</sub> assignments investigated herein may display long alkyl chains, which in turn produce  
507 daughter ions which undergo further fragmentation, leaving our analytical window at  $m/z > 150$ .

#### 508 *4.8. Implications for atmospheric carbon dioxide removal*

509 To meet the challenges of climate predictions at the end of the 21<sup>st</sup> century (IPCC, 2018),  
510 technologies will be needed that remove carbon dioxide from the atmosphere and sequester it for  
511 geological time frames. Such processes, termed carbon dioxide removal (CDR) are increasingly  
512 attracting both research and commercial investment interest. Only a few CDR technologies are  
513 currently being commercially developed, e.g. bio-energy with carbon capture and storage, or  
514 direct air capture, but are still far below the scale needed for globally significant CO<sub>2</sub> reduction,  
515 which would require achieving carbon drawdowns in the order of 100 to 1000 Gt by the year  
516 2100 (IPCC, 2018). A better understanding of natural mechanisms for long term preservation of  
517 organic molecules can provide guidelines for the development of technologies that could  
518 leverage those mechanisms to artificially enhance the removal of carbon in the geosphere in the  
519 form of engineered organic species. We term these species “alternative vectors for carbon  
520 storage”, or AVECS for short, and are currently investigating several routes to produce AVECS  
521 from organic precursors, including pathways involving sulfurization of organic molecules (Yim,  
522 2019). To that end, the study reported herein provided multiple indications relevant to AVECS  
523 technology development issues such as:

- 524 • Natural precursors such as pigments or lipids can incorporate sulfur as free molecules;
- 525 • Double bonds are the key molecular site for sulfur incorporation;
- 526 • Given the fact that double bonds are quite reactive, the sulfur incorporation must be  
527 taking place early in the diagenesis, at low pressures and temperatures, which favorably  
528 implies fewer energy needs for AVECS production;
- 529 • Cyclization is the favored sulfur incorporation mechanism producing thiolane structures,  
530 which can then be aromatized to thiophenes during diagenesis.

531 Some of these processes can be modelled and reproduced in the laboratory (Yim, 2019), however  
532 much more research effort will be needed to move these types of reactions from bench-scale to

533 sustainable pilot or large-scale deployments; notwithstanding, biogeochemical processes  
534 occurring in geosphere, such as those inferred from the current study of sulfur rich oils, can  
535 provide valuable insights and lessons which can direct this type of research.

## 536 **5. Conclusions**

537 The present study leveraged the expanded analytical window offered by novel analytical  
538 technologies to improve the current geochemical understanding of diagenetic processes  
539 involving sulfur and the resultant occurrence of sulfur rich oils. A set of five sulfur rich oils was  
540 analyzed via APPI-P FTICR-MS(/MS) and the following conclusions were drawn from the  
541 results:

542 The absence of peaks at  $m/z > 1250$  in Rozel Point and Jiangnan oils, as well as in their non-  
543 polar macromolecular fractions, disproves the previously established concept of a  
544 macromolecular network linked by (poly)sulfide bridges, which has been proposed to explain the  
545 occurrence of non-GC-MS amenable sulfur compounds in immature, sulfur rich oils. Instead, it  
546 appears that the presence of highly sulfurized lipids (up to six sulfur atoms) as free molecules are  
547 the result of sulfurization that in turn depends on the abundance of functional groups (double  
548 bonds mainly) in the precursor molecules.

549 Source rock thermal stress levels were found to be a key factor differentiating the composition of  
550 biodegraded, sulfur rich Peace River oils, compared to thermally immature, sulfur rich Rozel  
551 Point and Jiangnan oils. Whereas Peace River oils do not show biological precursor skeleton  
552 signatures, Rozel Point and Jiangnan oils show the presence of S-containing steroids, hopanoids,  
553 carotenoids and alkenones among others.

554 APPI-P FTICR-MS/MS experiments indicated only a minor occurrence of thiol functional  
555 groups and sulfide bridges linking two precursor molecules. The compositional patterns observed  
556 in sulfurized  $C_{40}$  assignments, representing S-bound carotenoids, supports a mechanism of sulfur  
557 incorporation involving the formation of thiolane structures, which can then be aromatized to  
558 thiophenes during diagenesis. FTICR-MS/MS experiments were also helpful to prove the  
559 occurrence of intermolecular sulfur incorporation reactions as a secondary mechanism, the  
560 aliphatic nature of class  $S_1$  DBE 1 species and sulfurized alkenones, the preference of cycle-

561 forming sulfur incorporation in steroid structures, and the occurrence of thiol/thioether groups,  
562 although these were seen only to a minor extent.

563 The analytical approach presented herein can be improved and expanded by using analytical  
564 workflows that include chromatographic separation to better characterize individual molecular  
565 species, as well as by including other sample types such as source rocks and recent sediments to  
566 better constrain the progression of diagenetic processes involving sulfur incorporation.

567 The occurrence of S-bound compounds in oils or sediment extracts, as detected via FTICR-MS,  
568 can be used as proxies for depositional environment, organic matter input, sulfurization and other  
569 diagenetic processes. Understanding the relationship between the precursor and sulfurized  
570 products, as well as the mechanisms for such reactions, is key to the design of geoengineering  
571 solutions to enhance carbon preservation in the geosphere by promoting sulfurization of biomass.

572

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### 576 **Notes**

577 The authors declare no competing financial interest.

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585 TBPO. CY carried out literature reviews, laboratory experiments, data interpretation and  
586 contributed to writing. FTICR-MS method optimization and analyses were carried out by MB,

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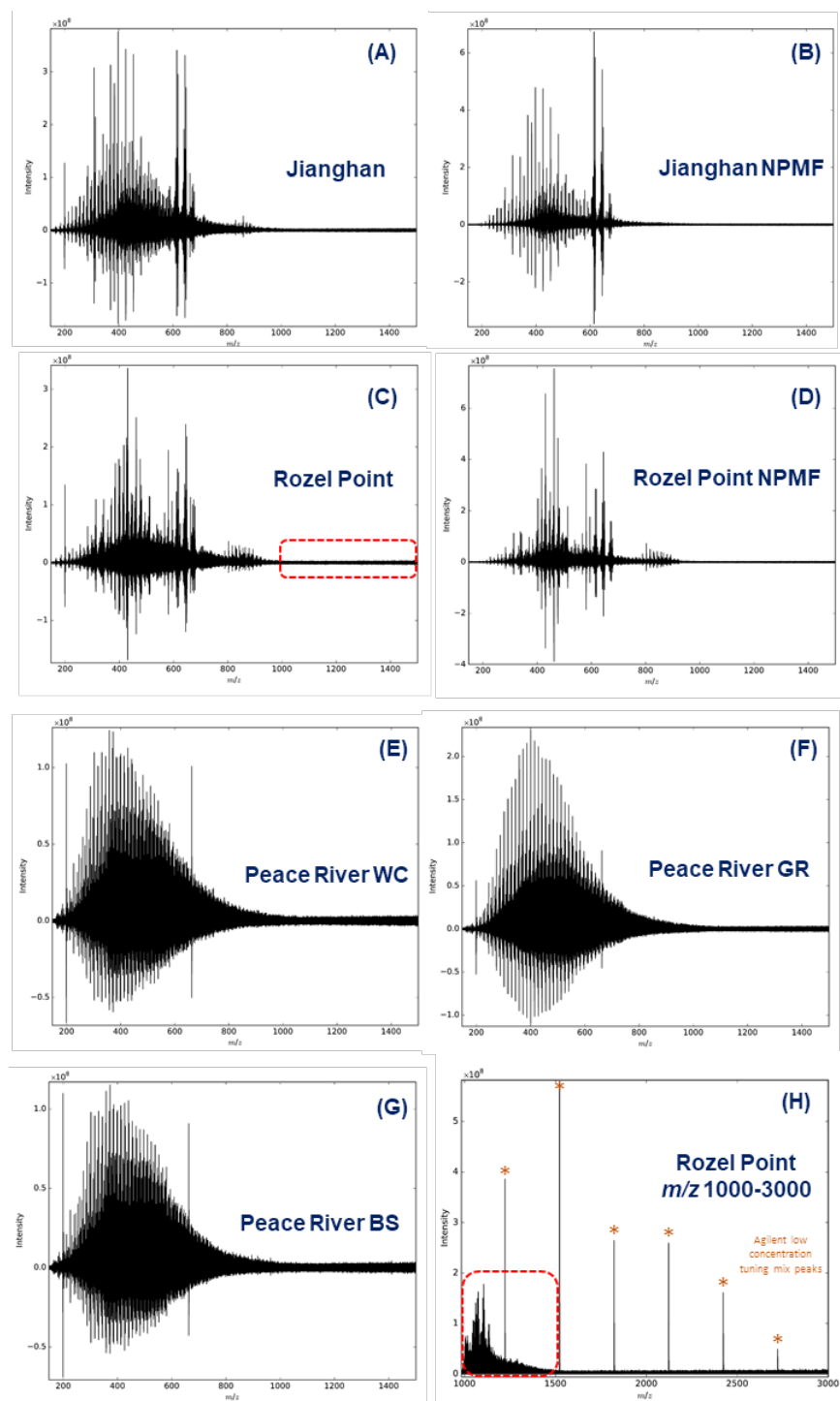
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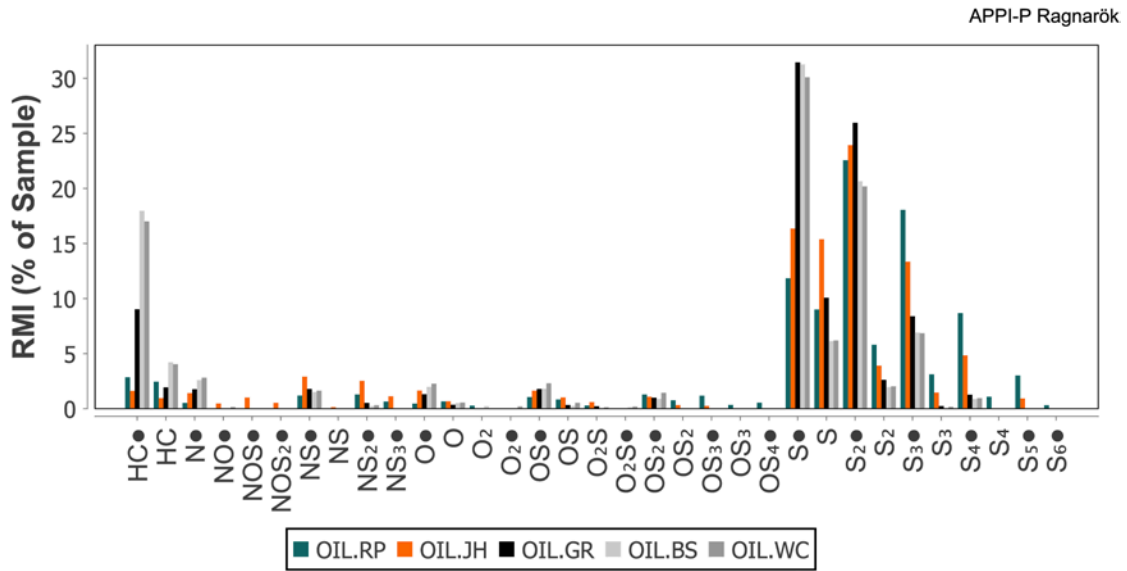
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1

2 **Figure 1.** APPI-P FTICR absorption-mode mass spectrum of (A-B) Jianghan oil and NPMF, (C-  
3 D) Rozel Point oil and NPMF, (E-G) Peace River oils (see Table 1 for sample codes). Plot H  
4 shows the Rozel Point oil APPI-P FTICR magnitude-mode mass spectrum, acquired from  $m/z$   
5 1000 – 3000, 100 scans, ion accumulation time increased by a 100-fold (500 ms), where the  
6 sample solution was doped with 2.5  $\mu\text{L}$  of the Agilent Low Concentration Tuning Mix.

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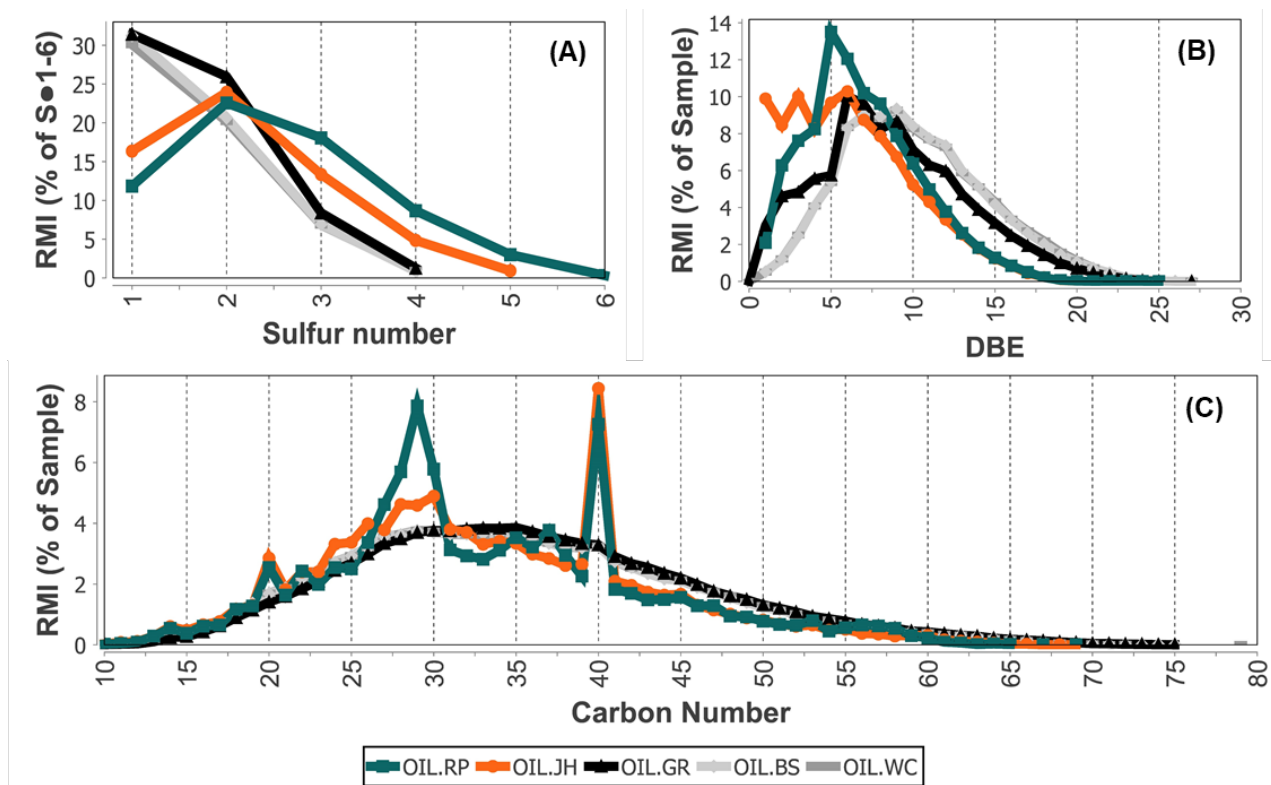
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9 **Figure 2.** APPI-P FTICR-MS compound class distribution of the oil samples analyzed in this  
10 study. Dots following class labels refer to radical ions (odd electron ion) while the absence of a  
11 dot refers to protonated ions (even electron ion). The relative monoisotopic intensity (RMI, % of  
12 Sample) is calculated as the fraction of the total monoisotopic intensities of a compound class,  
13 normalized to the sum of monoisotopic intensities of all compound classes in the sample. See  
14 Table 1 for sample abbreviations.

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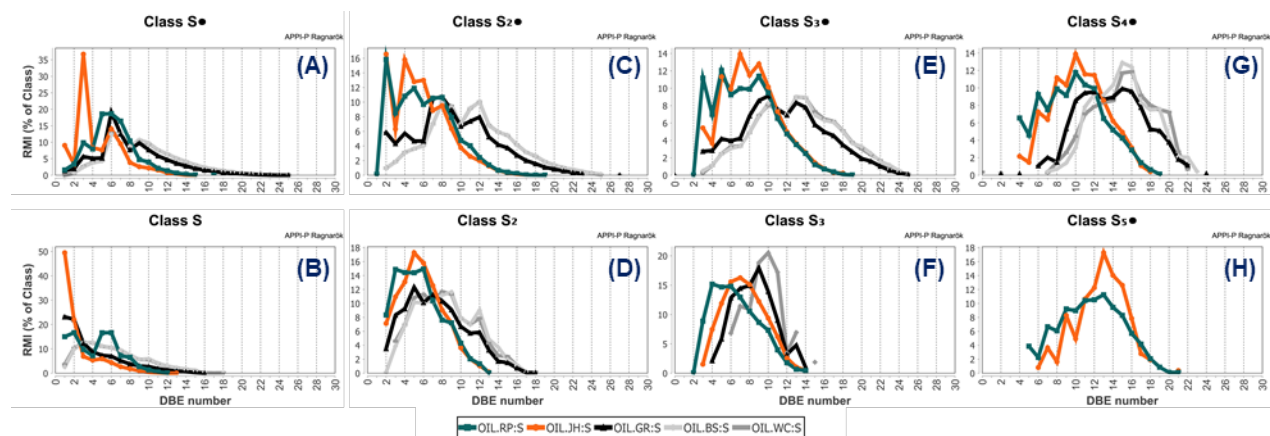
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19 **Figure 3.** APPI-P FTICR-MS parameter distributions in the oil samples analyzed in this study.  
20 (A) DBE number distribution of all assigned peaks; (B) Sulfur number distribution, based on  
21 radical heteroatomic classes  $S_{\bullet 1-6}$  peaks; (C) Carbon number distribution of all assigned peaks.  
22 See Table 1 for sample abbreviations.

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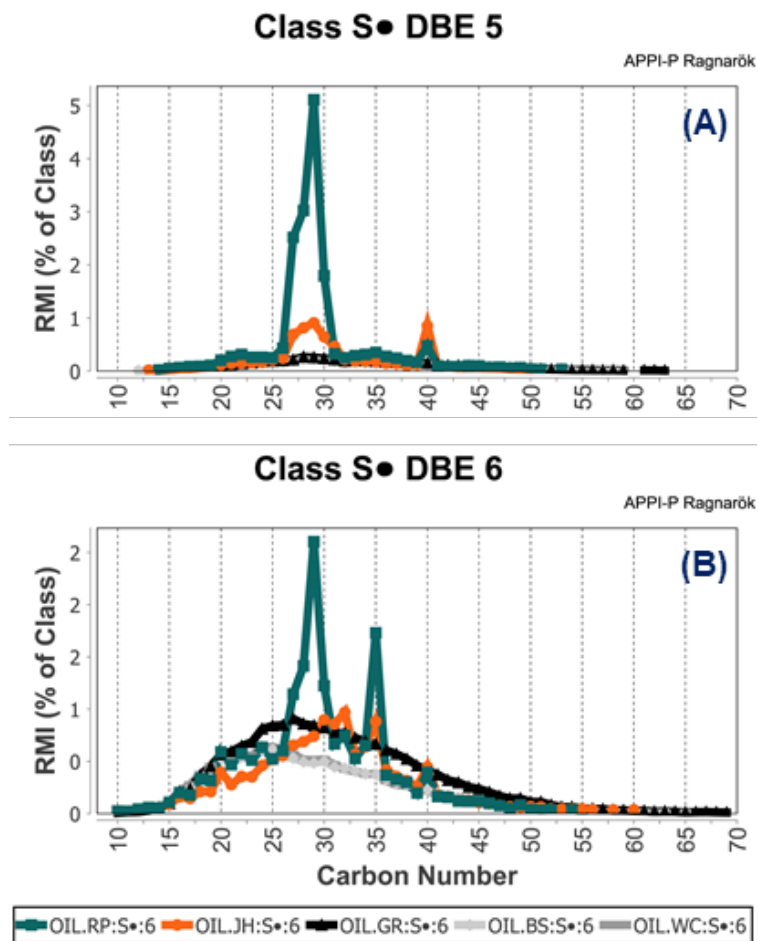
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27 **Figure 4.** APPI-P FTICR-MS heteroatomic classes  $S_{1-5}$  and  $S_{1-3}$  DBE number distributions in  
28 the oil sample set. Dots following class labels refer to radical ions (odd electron ion) while the  
29 absence of a dot refers to protonated ions (even electron ion). See Table 1 for sample  
30 abbreviations.

31



33

34 **Figure 5.** APPI-P FTICR-MS carbon number distribution of heteroatomic classes S•<sub>1</sub> DBE 5 (A)  
 35 and 6 (B). Dots following class labels refer to radical ions (odd electron ion. See Table 1 for  
 36 sample abbreviations.

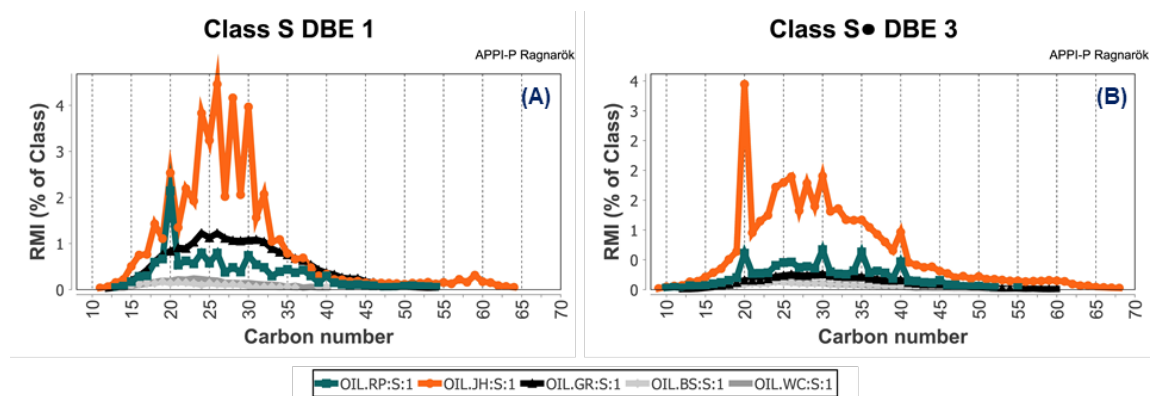
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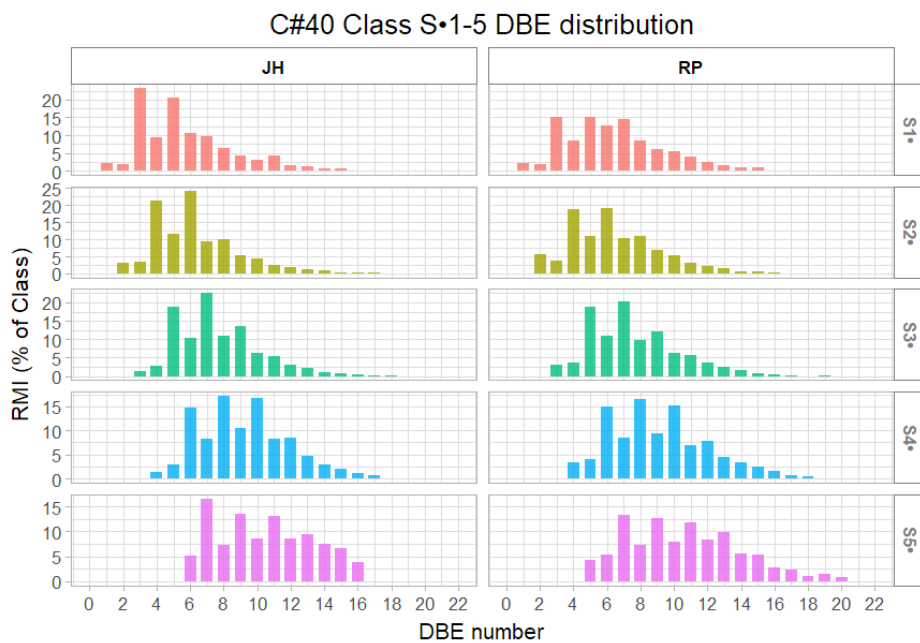
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42 **Figure 6.** APPI-P FTICR-MS carbon number distribution of heteroatomic classes (A)  $S_1$  DBE 1  
43 and (B)  $S_{\bullet 1}$  DBE 3. Dots following class labels refer to radical ions (odd electron ion) while the  
44 absence of a dot refers to protonated ions (even electron ion). See Table 1 for sample  
45 abbreviations.

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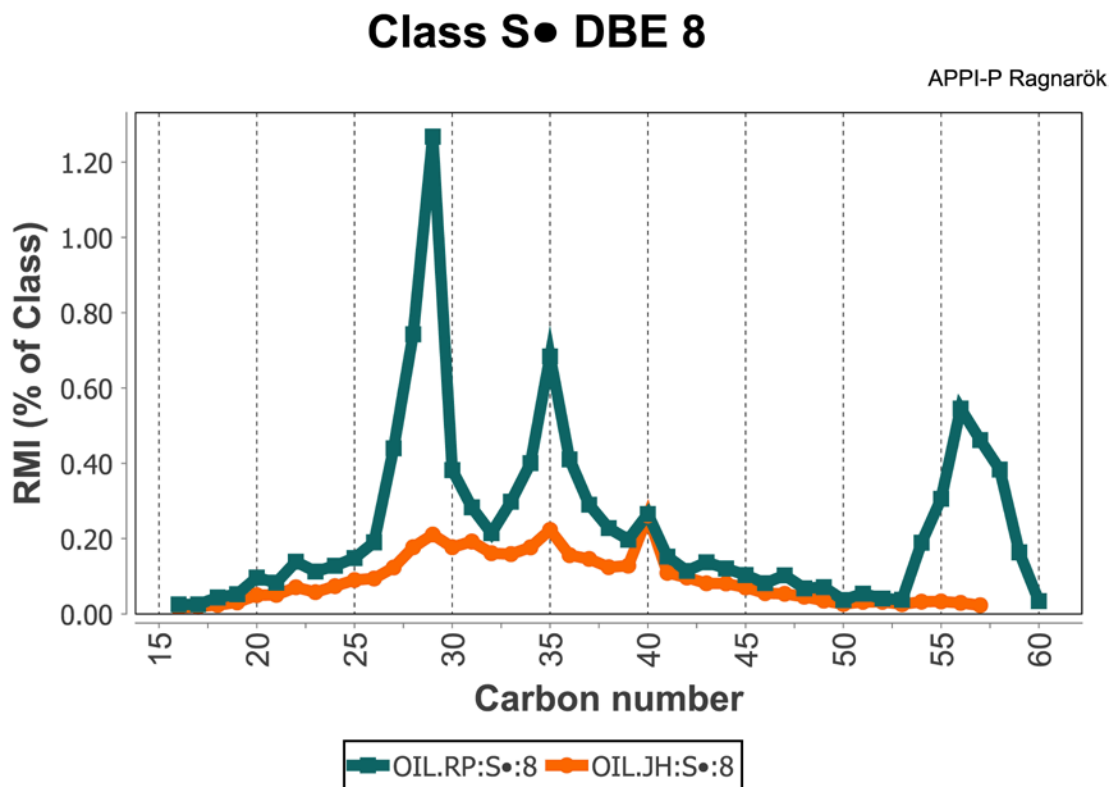
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50 **Figure 7.** DBE distribution of class S<sub>1-5</sub>• C<sub>40</sub> species in Rozel Point and Jiangnan oils, as  
51 detected by APPI-P FTICR-MS.

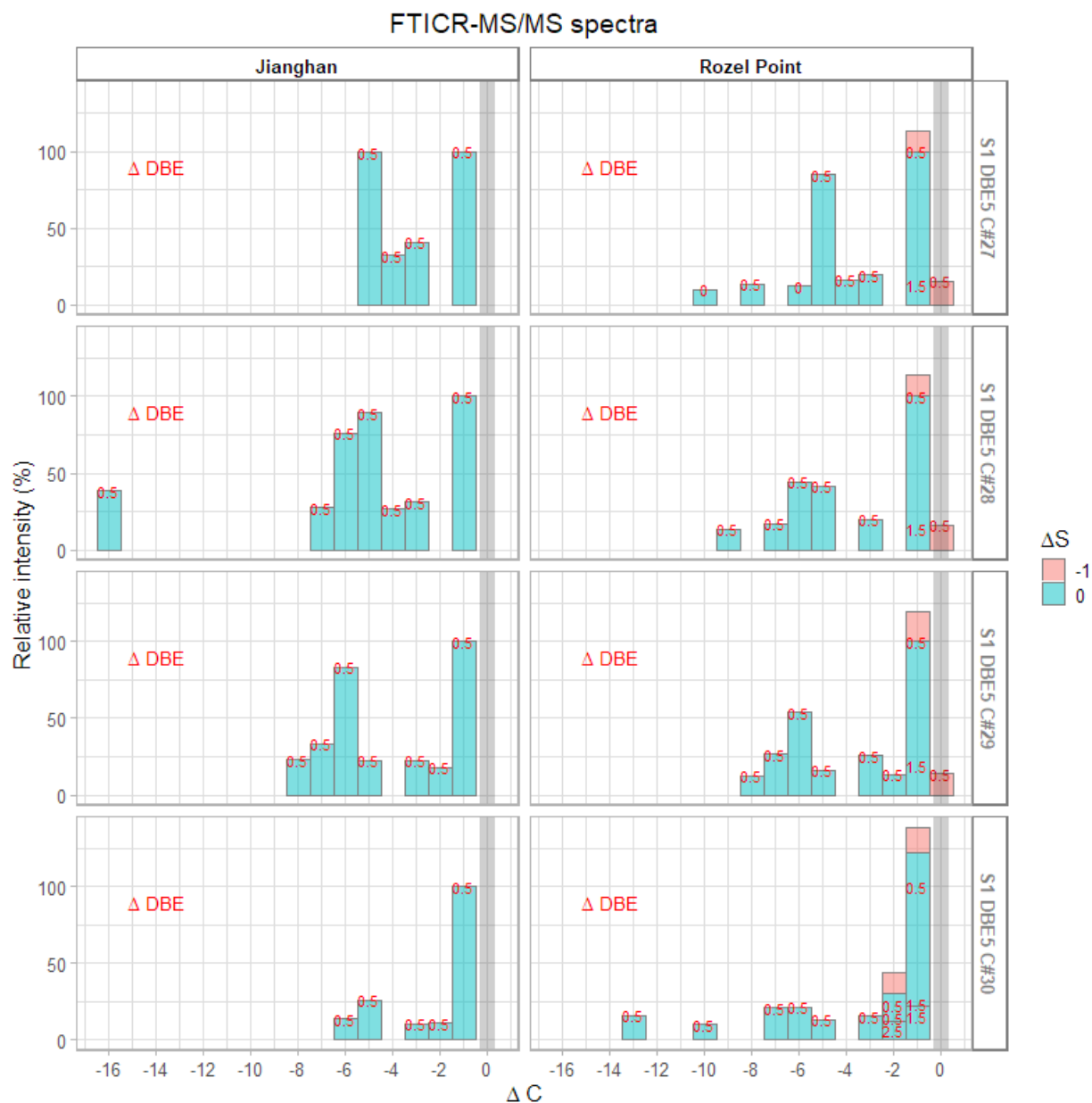
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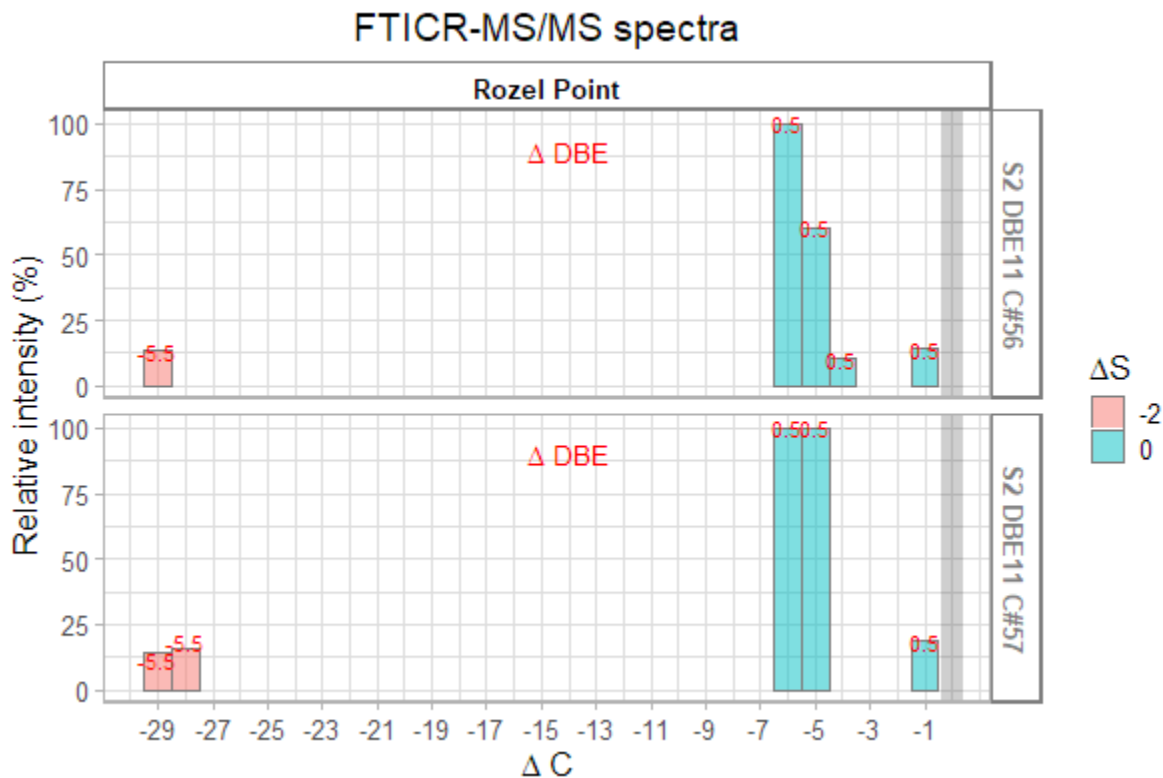
55 **Figure 8.** APPI-P FTICR-MS carbon number distribution of heteroatomic class S•<sub>1</sub> DBE 8, in  
 56 Jiangnan (JH) and Rozel Point (RP) oils. Dots following class labels refer to radical ions (odd  
 57 electron ion).

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59

60 **Figure 9.** Representation of the APPI-P FTICR-MS/MS spectra of class  $S_1$  DBE 5  $C_{27-30}$   
 61 species. Detected fragments are shown as bars based on their compositional differences to the  
 62 parent peak: carbon number (x-axis), sulfur number (colors) and DBE (labels). Relative  
 63 intensities are normalized to the most intense daughter ion. The parent peak, represented by a  
 64 gray bar with arbitrary y-axis values, is not considered in the relative intensity calculations.  
 65 Fragments with carbon or sulfur number higher than the parent peak, as well as those with  
 66 relative intensity <10%, are not shown.



67

68 **Figure 10.** Representation of the Rozel Point oil APPI-P FTICR-MS/MS spectra of class  $S_2$ •

69 DBE 11  $C_{56-57}$  species. Detected fragments are shown as bars based on their compositional

70 differences to the parent peak: carbon number (x-axis), sulfur number (colors) and DBE (labels).

71 Relative intensities are normalized to the most intense daughter ion. The parent peak, represented

72 by a gray bar with arbitrary y-axis values, is not considered in the relative intensity calculations.

73 Fragments with carbon or sulfur number higher than the parent peak, as well as those with

74 relative intensity <10%, are not shown.

75

76

1 **Table 1.** Sample codes and geochemical overview.

| <b>Code</b> | <b>Oil</b>  | <b>Location</b>             | <b>Sulfur (wt %)</b> | <b>Organic matter source</b>                 | <b>Depositional environment</b> | <b>Thermal maturity</b> | <b>Biodegradation</b> | <b>References</b>  |
|-------------|-------------|-----------------------------|----------------------|--|---------------------------------|-------------------------|-----------------------|--|
| RP          | Rozel Point | Utah, USA                   | 9.4                  | Marine - lacustrine                          | hypersaline, lacustrine         | Immature                | Moderate              | (Meissner et al., 1984; ten Haven et al., 1988; Sinninghe Damsté et al., 1989) |
| JH          | Jianghan    | Eastern China               | 5.6                  | Marine - lacustrine, with terrigenous inputs | hypersaline, lacustrine         | Immature                | Moderate              | (Philp and Zhaoan, 1987; Carroll and Bohacs, 2001; Hou et al., 2017)           |
| GR          | Peace River | NW- Central Alberta, Canada | 5.5                  | Marine - lacustrine, with terrigenous inputs | Marine - lacustrine             | Mature - overmature     | High                  | (Riediger, 1994; Adams et al., 2012; Adams et al., 2013; Bennett et al., 2013) |
| BS          | Peace River | NW- Central Alberta, Canada | 4.1                  | Marine - lacustrine                          | Marine - lacustrine             | Mature - overmature     | Very High             |  |
| WC          | Peace River | NW- Central Alberta, Canada | 4.2                  | Marine - lacustrine                          | Marine - lacustrine             | Mature - overmature     | High                  |  |

2