

What Role Does Hydrological Science Play in the Age of Machine Learning?

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Key Points:

- Hydrology lacks scale-relevant theories but deep learning experiments suggests that these theories should exist
- It is up to hydrologists to clearly show where and when hydrological theory adds value to simulation and forecasting

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Abstract

This paper is derived from a keynote talk given at the Google’s 2020 Flood Forecasting Meets Machine Learning Workshop. Recent experiments applying deep learning to rainfall-runoff simulation indicate that there is significantly more information in large-scale hydrological data sets than hydrologists have been able to translate into theory or models. While there is growing interest in machine learning in the hydrological sciences community, in many ways our community still holds deeply subjective and non-evidence-based preferences for models based on a certain type of ‘process understanding’ that has historically not translated into accurate theory, models, or predictions. This commentary is a call to action for the hydrology community to focus on developing a quantitative understanding of where and when hydrological process understanding is valuable in a modeling discipline increasingly dominated by machine learning. We offer some potential perspectives and preliminary examples about how this might be accomplished.

1 Beven’s Clouds

On April 27, 1900 William Thomson (Lord Kelvin) gave his ‘Two Clouds’ speech (‘Nineteenth-Century Clouds over the Dynamical Theory of Heat and Light’) at the Royal Institution, in which he argued that *“The beauty and clearness of the dynamical theory, which asserts heat and light to be modes of motion, is at present obscured by two clouds.”* The two open problems in physics that Kelvin referred to were the failure of the Michelson-Morley experiment to detect the luminous ether (*“how could the earth move through an elastic solid, such as essentially is the luminiferous ether?”*), and the ultraviolet paradox (*“the Maxwell-Boltzmann doctrine regarding the partition of energy”*). Within a decade, Einstein had proposed fundamentally novel insights that led to two paradigm shifts that define modern physics to this day - the transformation of these two ‘clouds’ into relativity and quantum mechanics.

In 1987, Keith Beven gave what might be considered hydrology’s version of the Two Clouds speech at a symposium of the International Association of Hydrological Sciences (IAHS) (Beven, 1987). He took a perspective inspired by Thomas Kuhn’s theory of scientific revolutions (Kuhn, 1962) to argue that *“[t]he extension of laboratory scale theory to the catchment scale is unjustified and that a radical change in theoretical structure (a new paradigm) will be required before any major advance can be made in [predicting catchment-scale rainfall-runoff responses].”* He proposed that two things would be necessary to push the field of surface hydrology into a new period of ‘normal science’: (i) scale-relevant theories of watersheds (*“[h]ydrology in the future will require a macroscale theory that deals explicitly with the problems posed by spatial integration of heterogeneous nonlinear interacting processes”*), and (ii) uncertainty quantification (*“[s]uch a theory will be inherently stochastic and will deal with the value of observations and qualitative knowledge in reducing predictive uncertainty.”*)

Unfortunately, hydrology has not had its Einstein (with all due respect to A. Einstein, 1926; H. A. Einstein, 1950). Nine decades from the establishment of the Hydrology section of the American Geophysical Union and after more than a half-century of computer-based hydrological modeling (Crawford & Burges, 2004), Blöschl et al. (2019) listed as one of the twenty three ‘Unsolved Problems in Hydrology’: *“what are the hydrologic laws at the catchment scale and how do they change with scale?”*

2 Tilting at Windmills

There are several potential reasons why the search for scale-relevant theories in hydrology has been unsuccessful, but lack of effort is not one of them (e.g., Beven, 2006b; Blöschl & Sivapalan, 1995; Dooge, 1986; Peters-Lidard et al., 2017; Sivapalan, 2006). One potential reason is simply that there might be no scale-relevant theories to find - it is

66 possible that macroscale watershed behaviors are dominated by heterogeneity, meaning
 67 that there is little consistency across different basins. As summarized by Hrachowitz et
 68 al. (2013), *“Beven (2000) highlighted the varying importance of different hydrological pro-
 69 cesses, active at different time scales in different catchments, and thereby emphasized unique-
 70 ness of place as a consequence of the variability of nature.”*

71 Alternatively, it could be the case that there are consistent macroscale patterns in
 72 hydrologic behaviors across watersheds, but we lack sufficient observations (type, scale,
 73 scope) to discover these similarities. Again, as summarized by Hrachowitz et al. (2013),
 74 *“[i]t was realized that increased physical model realism (and complexity) requires both more
 75 input data and more model parameters, which are rarely available with sufficient detail
 76 to account for catchment heterogeneity at the required resolution.”*

77 Uniqueness of place and lack of data are, in our experience, two of the most com-
 78 mon hypotheses about why hydrology lacks both scale-relevant theories of watersheds.
 79 The alternative to such hypotheses is that these theories could exist and that there is
 80 enough information in available observation data that we could have discovered them,
 81 but that hydrologists simply have failed to do so. Prior to last year, it is fair to say that
 82 as a community we did not know which of these reasons was the cause of our lack of suc-
 83 cess. However, with the accelerating development of modern machine learning (ML), and
 84 deep learning (DL) in particular, we know that the reason is the third one listed: watershed-
 85 scale theories (and models) could have been derived from currently-available observa-
 86 tion data, but the hydrology community simply failed to do so.

87 The reason that we know this is because general models can be learned with DL.
 88 In a large sample study using 30 years of data from several hundred basins in the con-
 89 tinental United States, DL gave better daily streamflow predictions on average in *un-
 90 gauged* basins than traditional hydrology models when calibrated to long data records
 91 in *gauged* basins (Kratzert, Klotz, Herrnegger, et al., 2019). That study used benchmarks
 92 based on (i) a modern process-based model that was the product of several millions of
 93 dollars of development funding, and (ii) a conceptual model calibrated separately for each
 94 individual basin (Figure 1). These DL models have been benchmarked against a num-
 95 ber of conceptual and process models calibrated both locally and regionally using a va-
 96 riety of metrics and hydrological signatures (Kratzert, Klotz, Hochreiter, & Nearing, 2020;
 97 Kratzert, Klotz, Shalev, et al., 2019). The fact that DL learned to predict in unseen basins
 98 better than traditional models in gauged basins indicates that there exists inter-basin
 99 consistency that we should be able to exploit and develop into a watershed-scale theory
 100 of rainfall-runoff behavior.

101 The problem of prediction in ungauged basins (PUB) (Hrachowitz et al., 2013; Siva-
 102 palan et al., 2003) is fundamentally a problem of extrapolation. Unlike both conceptual
 103 and process-based hydrology models, and also unlike shallow ML models that the hy-
 104 drological science community has used in the past (e.g., Hsu, Gupta, & Sorooshian, 1995),
 105 DL models work *better* when trained on multiple catchments than when trained on in-
 106 dividual catchments (Figure 2; also see the more thorough data scaling analysis by Gauch,
 107 Mai, and Lin (2019)). This means that DL models learn relationships from a large sam-
 108 ple of hydrological variability and are able to translate those learned relationships into
 109 better predictions in any individual basin. In contrast, traditional hydrology models are
 110 best when calibrated to individual basins, and performance always degrades when trans-
 111 ferring to other basins or when using regional calibration.

112 It is often claimed that one of the reasons hydrology models don’t extrapolate well
 113 is because they are over-calibrated. Hrachowitz et al. (2013) reported that *“several au-
 114 thors (Kirchner, 2006; McDonnell et al., 2007; Wagener, Sivapalan, Troch, & Woods,*

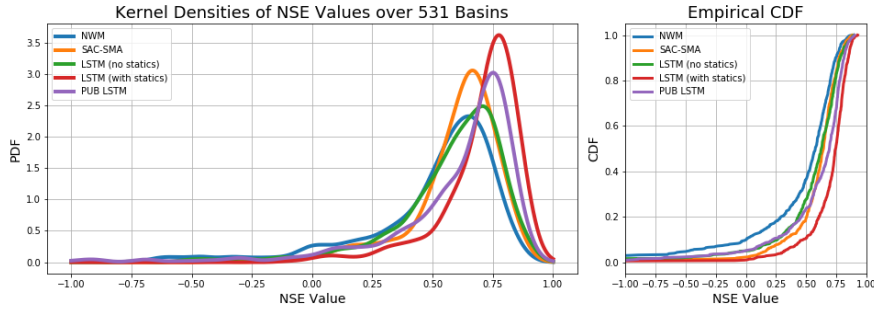


Figure 1. Results from Kratzert, Klotz, Herrnegger, et al. (2019) showing the empirical and cumulative distributions of model performance (Nash Sutcliffe Efficiencies) over a 15-year test period in 531 CAMELS catchments. SAC-SMA is the Sacramento Soil Moisture Accounting model, NWM is the National Water Model Reanalysis, and LSTMs are Long Short Term Memory networks (a type of deep learning architecture). The PUB-LSTM is the deep learning model applied in out-of-sample catchments. The other LSTM models (with and without statics) refer to deep learning models that were trained on all catchments (i.e., no out-of-sample catchments) and either did or did not use static catchment attributes (e.g., soils, vegetation, topography, etc.) as inputs.

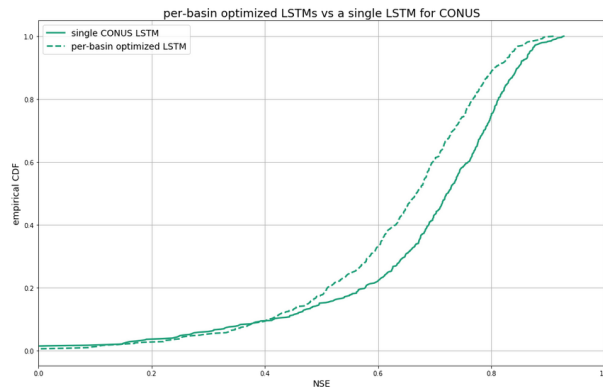


Figure 2. Cumulative distributions of NSE values over the same 531 CAMELS basins used by Kratzert, Klotz, Shalev, et al. (2019) from a single model trained over all basins (single CONUS LSTM) vs. separate models trained at each basin (per-basin optimized LSTM).

115 *2007)*¹ expanded on and strongly reiterated Klemeš's (1986b) arguments that models which
 116 perform adequately well during calibration, but fail to predict the hydrological catchment
 117 response in validation, frequently do so because they do not sufficiently represent the real-
 118 world processes that control the catchment response. Rather, their often high number of
 119 parameters together with the limited number of constraints (including both calibration ob-
 120 jectives and calibration criteria) resulted in high degrees of freedom, i.e. poorly condi-
 121 tioned parameter estimation problems, so that models behaved more like "mathematical
 122 marionettes." The problem is that this is not true. DL models generally have several or-
 123 ders of magnitude more degrees of freedom than calibrated conceptual models, and it
 124 is this *lack* of regularization that allows them to learn general and transferable hydro-

¹ These references are apparently incorrect in the quoted manuscript.

125 logical relationships. Using DL as a benchmark demonstrates that it is the regulariza-
 126 tion in the traditional models (i.e., the hydrological theory that the model structures are
 127 based on) that is actually the cause of their lack of generality and transferability, rather
 128 than this being a problem of over-parameterization.

129 To summarize, this benchmarking between DL and traditional hydrology models
 130 demonstrates three things. First that hydrologists could have developed general, scale-
 131 relevant theories of watersheds from available data, but failed to do so. Second, that our
 132 understanding of why such theories don't exist were incorrect - neither uniqueness of place
 133 nor lack of data was a valid reason for this failure. Third, that our understanding of why
 134 our existing models perform poorly in extrapolation is also incorrect - this is not due to
 135 a lack of regularization or to over-parameterization, but instead due to bad theory - the
 136 regularization (structure) that does exist in these models actively hurts us.

137 **3 Black Swans and Black Boxes**

138 In the preceding section, we argued that DL experiments suggests that new watershed-
 139 scale rainfall-runoff theory should exist, however DL models do not currently give us those
 140 theories. There are two ways we might think about this issue - both are currently open
 141 problems in hydrology.

142 First, we can leverage advances in explainable AI (XAI; Samek, 2019). It is often
 143 said that machine learning is a black box, and while there is some sense in which this
 144 is true, there is a much more important sense in which we should think about DL mod-
 145 els as containing complex, multi-layered, structured information that is accessible to us
 146 if we choose to query it. Recognizing this, our job as scientists becomes a problem of trans-
 147 lation: the information we want is in the models and we must learn how to translate that
 148 information into something that is human-interpretable.

149 Trained DL models typically don't yield new theory directly, however process-based
 150 models don't either. New insights from modeling studies come from probing models with
 151 various types of diagnostic tools (e.g., Martinez & Gupta, 2010; Nearing, Ruddell, Clark,
 152 Nijssen, & Peters-Lidard, 2018; Ruddell, Drewry, & Nearing, 2019; Yilmaz, Gupta, &
 153 Wagener, 2008), many of which are equally applicable to DL models. Examples of these
 154 tools are things like sensitivity analyses to understand (e.g.,spatiotemporal) input con-
 155 tributions (e.g., Sundararajan, Taly, & Yan, 2017), counterfactuals to understand cause
 156 and effect, (e.g., Pearl, 2013; Ribeiro, Singh, & Guestrin, 2016), or DL-specific tools like
 157 embedding layers and feature layer analyses (e.g., Bianchi, Rossiello, Costabello, Palmonari,
 158 & Minervini, 2020; Q. Wang, Mao, Wang, & Guo, 2017). We will give examples of hydrologically-
 159 relevant XAI in the context of the experiments described in Section 2 presently.

160 Second, we could use DL model for hypothesis testing. Instead of extracting in-
 161 formation from trained DL models, we can put hydrological theory into these models and
 162 assess improvement (or otherwise). From an ML perspective this is a regularization prob-
 163 lem, and common methods include things like (i) regularizing the loss function to pe-
 164 nalize violations of physical principles like conservation, monotonicity, etc. (e.g., Nabian
 165 & Meidani, 2020), (ii) augmenting scientific models with DL structures (e.g., Pelissier,
 166 Frame, & Nearing, 2020; Rackauckas et al., 2020) and (iii) architecturally constrained
 167 neural networks (e.g., Beucler et al., 2019; Daw et al., 2020).

168 According to most interpretations of the scientific method, hypotheses are tested
 169 by comparing predictions with observations. The results discussed in Section 2 can be
 170 interpreted as a hypothesis test that compares the information content of hydrological
 171 theory as encoded into models relative to a null hypothesis derived from data (Nearing,
 172 Ruddell, Bennett, Prieto, & Gupta, 2020). This does not mean that all hydrological the-
 173 ory encoded in those models should be rejected, but it is a challenge to disaggregate the
 174 good parts of that body of theory - that may provide significant hydrological informa-

175 tion - from the bad (H. V. Gupta, Wagener, & Liu, 2008). This is a classic problem of
 176 *underdetermination* (Laudan, 1990). In this case, the problem is due in part to the fact
 177 that for complex systems like watersheds it is necessary to aggregate a large number of
 178 different theories and bridge principles (Nagel, 1961) into predictive models. From a philo-
 179 sophical (and completely untested) perspective, we suggest that DL might help with the
 180 underdetermination problem to some extent by providing a modeling framework that
 181 allows us to aggregate pieces of hydrological theory into a functional, integrated model
 182 without requiring that the model includes theories of everything. For example, we might
 183 test hypotheses about conservation at various scales (or closure in data) without requir-
 184 ing an explicit assumption about infiltration, evapotranspiration, or groundwater sim-
 185 ply by embedding conservation laws into DL models (e.g., see Section 7.2). This is not
 186 possible with purely process-driven models that need descriptions of everything in or-
 187 der to account for every relevant catchment process. DL can learn the basic functional
 188 relationships from data and we can, in principle, assess the information content of any
 189 particular hypothesis by adding that hypothesis as a constraint on the DL model. While
 190 this paper was in review, an excellent example of this was provided by (Jiang, Zheng,
 191 & Solomatine, 2020), who included process modules as layers in deep tensor networks.

192 An example of the looking for explainability in a trained model is in Fig. 3. This
 193 figure shows the sensitivity of a time series DL model to past inputs. The model learned
 194 to store winter precipitation and release this as runoff when temperature and radiation
 195 increased in the spring. (Kratzert, Herrnegger, Klotz, Hochreiter, & Klambauer, 2019)
 196 showed that a DL time series model trained with inputs of precipitation and daily air
 197 temperature and targets of only daily streamflow contained internal states that corre-
 198 lated with snow cover and soil water storage. They showed that these ‘snow’ states were
 199 sensitive to inputs only when temperatures were below zero. None of this behavior was
 200 prescribed a priori - the model learned hydrologically-relevant, interpretable behavior
 201 about latent (unobserved) variables.

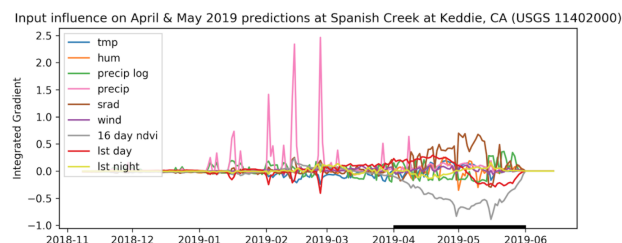


Figure 3. Sensitivity analysis using integrated gradients (Sundararajan et al., 2017) that shows the relative contributions to simulated streamflow during the months of April-May (heavy black shading on the x-axis) from the time-series of past inputs. The DL model learns to store winter precipitation and responds to increasing temperature and solar radiation in the spring.

202 Looking at the transferability and catchment similarity issues discussed in Section
 203 2, Kratzert, Klotz, Shalev, et al. (2019) constructed a DL network with an embedded
 204 feature layer that quantified catchment similarity along a number of learned dimensions
 205 (Figure 4). The features extracted from the trained network represent how the DL model
 206 transformed observable catchment characteristics into a representation of similarity and
 207 diversity in rainfall-runoff relationships. This matrix looks a little like noise, but it is a
 208 better representation of catchment similarity than anything human scientists have so far
 209 been able to develop. If we want to understand the information encoded in this matrix,
 210 then the job ahead of us is to translate this information into a human-interpretable form.
 211 Kratzert, Klotz, Shalev, et al. (2019) used dimensionality reduction to relate first-order

212 features in this similarity matrix with observable catchment characteristics and found
 213 that vegetation type and seasonality were the dominant influences.

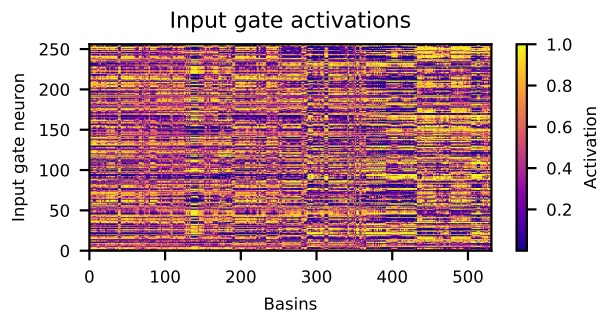


Figure 4. Results from Kratzert, Klotz, Shalev, et al. (2019) showing a matrix representing catchment similarity as identified by a deep learning model. There are 531 catchments (x-axis) and 256 model states (y-axis). Each state is activated for any individual catchment to some degree in the range $[0,1]$, with 0 meaning that the state is not used for that particular catchment. Similar catchments share more of this state space and dissimilar catchments share less.

214 While ML has been used in hydrology for decades, the ability (at least partially
 215 due to computational advances) to arrange shallow learning models into complex struc-
 216 tures with feature layers that can learn multi-scale patterns opens the door to leverag-
 217 ing diverse (e.g., multi-catchment) data in interpretable ways. The idea that ML mod-
 218 els are ‘black boxes’ is more of a testament to a lack of inspection, rather than to a fun-
 219 damental limitation of the models themselves. It’s worth noting that the DL models used
 220 by Kratzert, Klotz, Shalev, et al. (2019) were invented around the same time (Hochre-
 221 iter, 1991; Hochreiter & Schmidhuber, 1997) as some of the earliest shallow neural net-
 222 work applications in hydrology (e.g., Hsu et al., 1995). As a discipline, we have not done
 223 a great job of keeping pace with developments in ML.

224 4 Known Unknowns

225 The second ‘cloud’ in Beven’s (1987) speech was uncertainty. There has been an
 226 enormous amount of attention paid to this topic in the hydrological sciences (e.g., Beven,
 227 2006a, 2009, 2016; Beven & Binley, 2014; Beven, Smith, & Freer, 2007; Beven, Smith,
 228 Westerberg, & Freer, 2012; Beven, Smith, & Wood, 2011; Beven, Smith, & Freer, 2008;
 229 Clark, Kavetski, & Fenicia, 2011; P. Kumar, 2011; Mantovan & Todini, 2006; Montanari,
 230 2007; Montanari & Koutsoyiannis, 2012; Nearing, 2014; Nearing & Gupta, 2018; Pap-
 231 penberger & Beven, 2006; Renard, Kavetski, Kuczera, Thyer, & Franks, 2010; Stedinger,
 232 Vogel, Lee, & Batchelder, 2008; Todini & Mantovan, 2007; Vrugt, Ter Braak, Gupta, &
 233 Robinson, 2009), however we have not had a major breakthrough that led to a paradigm
 234 shift. We’ve suggested previously (Nearing, Tian, et al., 2016) that the uncertainty lit-
 235 erature in hydrology is somewhat detached from the discussion about uncertainty that
 236 is taking place in the larger academic (science and philosophy) communities. However,
 237 irrespective of that opinion, our community has not developed the stochastic theory of
 238 watersheds that Beven (1987) anticipated.

239 Dooge (1986) offered a discussion about why finding scale-relevant laws is difficult
 240 in many branches of science. His argument was that there are two basic categories of sci-
 241 entific theory: mechanistic and aggregate. In the former - mechanistic theories - we track
 242 properties (e.g., position, velocity) of individual components of a system, and the result-

243 ing model is usually expressed as a system of partial differential equations (PDEs). In
 244 the latter - aggregate theories - we rely on ergodic properties like the law of large num-
 245 bers to derive consistent statistical approximations (e.g., temperature, density) at scales
 246 that are much larger than the individual components of a system. The prototypical ex-
 247 ample of a mechanistic-type theory are Newton's laws, and the prototypical example of
 248 an aggregate-type theory is thermodynamics. Dooge borrowed the image in Figure 5 from
 249 Weinberg (1975) to illustrate this dichotomy - watersheds live in the middle area of or-
 250 ganized complexity, where complexity (heterogeneity) is at a similar scale to random-
 251 ness (lack of information).

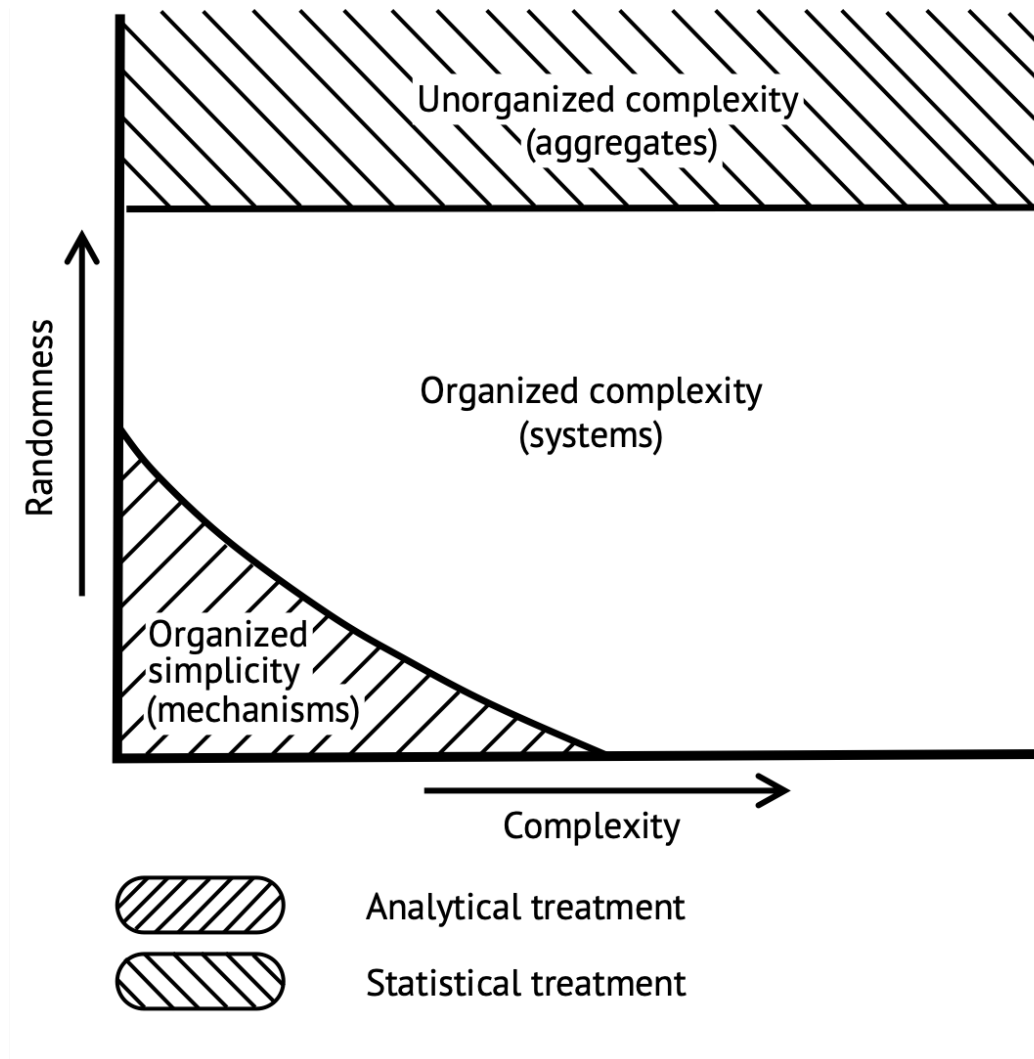


Figure 5. Recreation of an illustration that Dooge (1986) borrowed from Weinberg (1975) to show different types of successful theories in science. Watersheds arguably live in the area of organized complexity, where the complexity (heterogeneity) is at a similar scale to the randomness (lack of information).

252 Even imagined a hydrological theory that is fundamentally stochastic to account
 253 for heterogeneity. This is different than how hydrologists currently treat uncertainty. Typ-
 254 ical modeling approaches are mechanistic and treat a lack of complete information by
 255 adding additional (usually probabilistic) structure to a modeling problem. What we mean

256 by this is that our basic hydrologic theories are largely deterministic, and we represent
 257 lack of complete information by adding distributions on top of model inputs, structures,
 258 and predictions. Intuitively, it seems odd that we add *more* structure to a problem to
 259 represent a lack of information. Beven’s (1987) view of hydrologic theory is compelling
 260 in the sense that it would be preferable to have a theory of watersheds that is itself an
 261 aggregate-type theory, since at least a significant portion of the variability and complex-
 262 ity in watershed behaviors are due to both landscape and process heterogeneity.

263 Machine learning offers something like this in a straightforward way. Instead of pre-
 264 dicting the quantities of interest directly, we can predict distributional representations
 265 (e.g., probabilistic, fuzzy, etc.) directly from input data. This can be as simple as hav-
 266 ing the output of a DL model be the parameters of a parametric distribution (e.g., a mix-
 267 ture density, Bishop, 1994), or the quantiles of a nonparametric distribution (Taylor, 2000).
 268 An example of this is shown in Figure 6, which shows the weights of a mixture density
 269 over streamflow predicted by a DL model. The training loss function in this case was
 270 a likelihood function, and the model did not learn the mixture density parameters di-
 271 rectly, instead it learned how to predict these parameters from dynamic inputs. This fig-
 272 ure shows that the individual kernels of the mixture density respond in hydrologically-
 273 relevant ways - for example, some of the mixture weights have a seasonal cycle, and some
 274 are active only in rising or falling limbs of the hydrograph. It is important to understand
 275 that the DL model here maps directly from inputs (atmospheric forcings and static basin
 276 attributes; Addor, Newman, Mizukami, & Clark, 2017) to predicted probabilities, rather
 277 than sampling a priori probabilities over different model components. There is no need
 278 to prescribe any a priori probabilities.

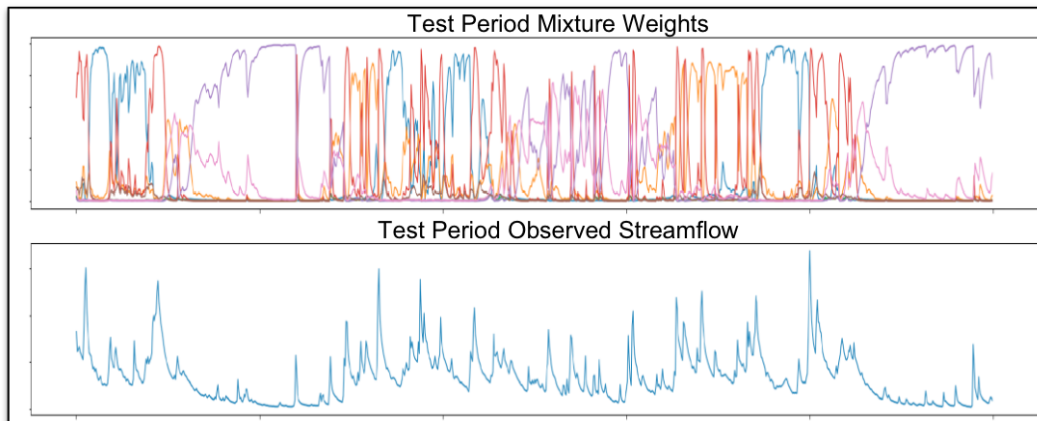


Figure 6. Mixture density weights (6 kernels) predicted by a deep learning model (top) as compared with the corresponding observed hydrograph (bottom). The mixture density weights vary in hydrologically-relevant ways - i.e., as a function of peaks (red) and recessions (blue).

279 We find an important distinction between *generative* vs. *discriminative* models (Near-
 280 ing, Gupta, & Crow, 2013). Generative models produce a joint distribution between tar-
 281 gets, Y , and inputs, X , and then invert that distribution to obtain conditional predic-
 282 tive probabilities $p(Y|X)$. Discriminative models, on the other hand, map directly onto
 283 conditional probabilities. Discriminative models avoid the need to assign any a priori prob-
 284 abilities, and if we believe that we have some information about uncertainties associated
 285 with various inputs, these uncertainties can always be used as additional inputs into the
 286 model.

287 Traditional hydrology models, on the other hand, are generative. We must first de-
 288 fine all input distributions and our predicted distributions come from sampling those a
 289 priori prescribed distributions. When we use an ensemble to represent uncertainty, for
 290 example, the hydrological model or family of models produces a joint distribution be-
 291 tween inputs and targets. Although we can sample the predictive conditional by sim-
 292 ply looking at one ensemble member, the distribution itself does not exist except as im-
 293 plied by the ensemble where each ensemble member is a joint sample of (X, Y) . The bot-
 294 tom line is that in a generative approach, the predicted probabilities are defined in ad-
 295 vance by the input or sampling probabilities.

296 While aggregate theories exist for certain hydrological fluxes (e.g., Singh, Yang, &
 297 Deng, 2003; J. Wang & Bras, 2011), most operational models are based on mechanis-
 298 tic theories - hydrologists have not developed an aggregate theory of watersheds. ML does
 299 not produce aggregate theories, but it does allow for discriminative modeling.

300 In addition to predicting probabilities directly, discriminative ML models can take
 301 *any type of input*, given sufficient training data. This offers an alternative to inverse meth-
 302 ods like data assimilation for integrating ancillary data streams (Nearing et al., 2013).
 303 Feng, Fang, and Shen (2019), for example, used the discriminative approach to integrate
 304 lagged streamflow values in a (deterministic) DL streamflow model. In principle, it is fea-
 305 sible to add any type of input into one of these models as long as there is sufficient train-
 306 ing data. We no longer need to prescribe the various input distributions directly - in-
 307 stead these are learned (either implicitly or explicitly) by the DL model from all avail-
 308 able data in a way that is dynamic (i.e., changes) in time and place, and under differ-
 309 ent hydrologic conditions.

310 5 Overlapping Magisteria: Faith and Fact in Hydrology

311 In the previous sections, we motivated several arguments highlighting conceptual
 312 deficiencies in hydrological science that were demonstrated by recent findings from bench-
 313 marking DL models. This type of benchmarking result is not new - hydrologists have
 314 been testing ML models against both calibrated conceptual models and process-based
 315 models for at least a quarter century (Hsu et al., 1995, depending on how we define ML),
 316 and it has always been the case that ML generally outperforms other types of models
 317 (e.g., Abramowitz, 2005; Best et al., 2015; Nearing, 2013; Nearing, Mocko, Peters-Lidard,
 318 Kumar, & Xia, 2016).

319 Todini (2007) framed the issue like this: *“physical process-oriented modellers have*
 320 *no confidence in the capabilities of data-driven models’ outputs with their heavy depen-*
 321 *dence on training sets, while the more system engineering-oriented modellers claim that*
 322 *data-driven models produce better forecasts than complex physically-based models.”* The
 323 key phrases in this sentence are ‘*confidence in*’ and ‘*better forecasts*’ - one is a statement
 324 of belief and one is a statement of fact.

325 Hydrology as an applied science is motivated by both *epistêmê* and *technê* (Parry,
 326 2003). On one hand (*technê*), we are often funded to tackle acute societal needs for man-
 327 aging water resources and and mitigating water-related hazards. On the other hand (*epistêmê*),
 328 many of us are true curiosity-driven scientists and care fundamentally about increasing
 329 our understanding of the world around us. These two objectives, however, cannot be cleanly
 330 separated. Whether any individual hydrologist is personally motivated by societal rel-
 331 evance vs. primal curiosity (the analogy we want the reader to draw is with Gould’s (1999)
 332 claim that *“science treats factual reality, while religion treats human morality”*), - the
 333 fact is that scientific hypotheses are tested by their ability to make accurate predictions.
 334 If our hypotheses do not translate into consistent accurate predictions, then they have
 335 not passed the basic test of science. The situation is a little more complicated when com-
 336 paring the information content of data-driven vs. theory-driven models, since imperfect

337 or incomplete theory can still be valid and useful, but hypotheses only become part of
 338 a body of theory if they translate into consistently accurate predictions.

339 The trend in the hydrology community has been toward more detailed process-based
 340 models based on essentially old theories of closure. As an example, Wood et al. (2011)
 341 suggested that “*developing a hyperresolution hydrological prediction capability is a “grand*
 342 *challenge for hydrology” because of the significant modeling, computational, and data needs*
 343 *that will be required for global or continental predictions at these spatial resolutions [~*
 344 *100m].’* This was cited as a major part of the motivation for developing the US Na-
 345 tional Water Model (Salas et al., 2018), which doesn’t out-perform simpler modeling strate-
 346 gies (e.g., Figure 1). Is the idea that if we keep increasing resolution and complexity, our
 347 models will reach a tipping point or there will be a step change in accuracy? Are we look-
 348 ing for incremental improvements with a trajectory sufficient to catch up to the accu-
 349 racy we get from ML, even as the pace of development in basic ML and AI science in-
 350 creases and the Earth-observation record available for training continues to grow? Is this
 351 a reasonable expectation that more of the same will help solve the fundamental prob-
 352 lem (lack of scale-relevant theory)?

353 In their report of the IAHS community-wide effort to outline key ‘Unsolved Prob-
 354 lems in Hydrology’ (UPH), Blöschl et al. (2019) said that “*[m]ost hydrologists would prob-*
 355 *ably agree that [extrapolating to changing conditions] will require a more process-based*
 356 *rather than a calibration-based approach as calibrated conceptual models do not usually*
 357 *extrapolate well.”* Similarly, in a summary of a recent workshop on ‘Big Data and the
 358 Earth Sciences’ Sellars (2018) reported that “*[m]any participants who have worked in*
 359 *modeling physical-based systems continue to raise caution about the lack of physical un-*
 360 *derstanding of ML methods that rely on data-driven approaches.”* The problem with these
 361 types of opinions is that in any case where we have sufficient observation data to bench-
 362 mark models, ML does better, even out-of-sample (see references above). Similarly, Kirchner
 363 (2006) claimed that “*[i]t is almost axiomatic that we need “physically based” mod-*
 364 *els in order to make reliable predictions beyond the range of prior observations.”* This
 365 is not an axiom of any theorem or any tautology, it is a hypothesis at best, and one that
 366 has failed every empirical test put to it that we are aware of. This is not science, it is
 367 religion.

368 The reason that DL in particular has at least the potential to remain reliable un-
 369 der changing conditions is because these models can be trained on a large diversity of
 370 data. As any particular catchment changes, it is likely that there are other catchments
 371 in a global data set that is similar along one or more of the changing dimensions. There
 372 will always be some catchments that evolve outside of the training envelope in terms of
 373 climate change or other anthropogenic influences, and it is unknown how model (of any
 374 type) will behave in such situations. But most catchments in the world will have some
 375 analogue along most dimensions of climate or land use, etc. It will be a critical project
 376 to understand how to structure the right mix of theory and data for developing reliable
 377 models at, for example, the climate scale, but the presumption that such projections must
 378 be “physically-based” seems strange. Why would we ever prefer a model that does worse
 379 on the data that we actually have in-hand?

380 6 Hydrology Beyond Streamflow

381 The hydrological sciences are diverse and the discussion so far has been about catch-
 382 ment hydrology and streamflow. Supposing the reader accepts the arguments we’ve laid
 383 out so far, it’s worth asking whether there are implications for other branches of the dis-
 384 cipline. The answer is - of course - that we don’t know. On one hand, there are major
 385 differences between the challenges faced in catchment hydrology vs. groundwater or eco-
 386 hydrology or hydrometeorology, but at the same time it is difficult to overestimate the
 387 impact of DL and AI throughout all types of human endeavors. In hydrometeorology sev-

388 eral studies have shown that even very simple regression models produce better estimates
 389 of radiation partitioning than process-based land surface models Abramowitz (2005); Best
 390 et al. (2015); Nearing, Ruddell, et al. (2018). Fang and Shen (2020) showed that DL can
 391 produce highly accurate soil moisture forecasts with remote sensing. Hydrometeorology
 392 is similar to streamflow hydrology in that observations are (relatively) abundant from
 393 satellites and mature sensor networks like FluxNet, etc. These fields are also similar in
 394 that the major sources of uncertainty are due to spatial heterogeneity at intermediate
 395 scales.

396 In groundwater, which is often more data limited than surface hydrology, many of
 397 the standard methods have close or direct analogs in ML already (e.g., Kriging is just
 398 Gaussian process regression Williams and Rasmussen (2006)). It may be the case that
 399 there is less potential for a fundamentally new result. One recent study reported that
 400 a physically-based groundwater model outperformed several shallow ML models (Chen,
 401 He, Zhou, Xue, & Zhu, 2020). There have been some relatively small DL studies in ground-
 402 water hydrology (e.g., Mo, Zabararas, Shi, & Wu, 2019; Sahoo, Russo, Elliott, & Foster,
 403 2017) that did not report transformative results.

404 It is hard to draw strong conclusions from the existing body of work. In all of these
 405 studies (including those by the current authors but with the notable exception of Fang,
 406 Pan, and Shen (2018)) is a lack of big data. ML does not have the ability to learn multi-
 407 scale hierarchical patterns in the same way as DL, and therefore cannot leverage diver-
 408 sity in big data in the same way. After testing several shallow ML models, Chen et al.
 409 (2020) concluded that *“the generalization ability of numerical model is superior to the
 410 machine learning models because of the inclusion of physical mechanism.”*

411 The basic problem is a lack of real investment into this type of effort. There are
 412 major programs across hydrologic disciplines to build comprehensive multi-scale mod-
 413 els (e.g., groundwater (de Graaf, Condon, & Maxwell, 2020), streamflow Li et al. (2015);
 414 Lin et al. (2019), hydrometeorology, (Rodell et al., 2004), and many others) but to our
 415 knowledge there is no similar effort to build global AI models. DL does not scale like tra-
 416 ditional models - it works differently on large data sets than small data sets, - so small
 417 pilot studies do not tell us much.

418 There is no question that we are in a new information age, and that modern data
 419 science techniques have been transformative across scientific disciplines. The message
 420 that we would like to leave the reader with is that hydrologists currently don’t know what
 421 how transformative this technology will across our discipline. We do not know this be-
 422 cause we have not made a serious investment in AI-based hydrology. Our major mod-
 423 eling centers continue to invest primarily in old technologies and old approaches. In the
 424 case of streamflow hydrology, this has been a disaster. The point of this opinion piece
 425 is that there are clues that maybe the balance of data and theory will not look like what
 426 hydrologists anticipate (e.g., references in Section 5).

427 7 Where the Sidewalk Ends

428 So what could we do about this? The following subsections outline what we see as
 429 both immediate needs for expanding DL in hydrology, as well as some ideas about what
 430 the longer-term future could look like.

431 7.1 Distributed Modeling

432 The first immediate need is for spatiotemporal DL models in all areas of hydrol-
 433 ogy. We simply just need to make serious investments across the discipline to gather the
 434 data that each community has - across regions and countries, to the extent possible - and
 435 make a serious attempt to develop state-of-the-art AI models.

436 We expect that first-order attempts at this type of project will look similar to cur-
 437 rent models with some explicit spatiotemporal extent/resolution and some number of
 438 latent (hidden) variables. Previously, we criticised calls for hyper-resolution modeling,
 439 and while the race to higher-density, more-of-the-same type models does seem to be a
 440 particularly unthoughtful idea, it is nevertheless the case that hydrological processes have
 441 both spatial and temporal components. We expect that within the next 1-2 years the
 442 community will develop several distributed DL watershed models (e.g., Moshe et al., 2020).

443 There are various ways that we might incorporate a multitude of different types
 444 of spatiotemporal data into trained models. DL allows for complex interactions between
 445 different feature layers, and fine tuning allows modelers to train individual components
 446 of a model. We can imagine a model developed by training different feature layers - per-
 447 haps themselves multi-layer DL models - and piecing these together to represent theory-
 448 guided architectures. As an example, we could imagine training a convolutional network
 449 to map from remote sensing data like SMAP (Entekhabi et al., 2010) to root-zone soil
 450 moisture by training directly on target data from in situ networks like the USDA Soil
 451 Climate Analysis Network (Schaefer, Cosh, & Jackson, 2007) and/or FluxNet (Baldoc-
 452 chi et al., 2001). The weights of this trained convolutional layer(s) could then be frozen,
 453 and the trained network then used as one (of many) input feature layer(s) into an LSTM
 454 (or other time series model) for predicting streamflow (or evapotranspiration or ground-
 455 water recharge). In principle, input data streams could be integrated at arbitrary spa-
 456 tiotemporal resolutions so that irregular convolutional networks (e.g., graph convolutions)
 457 could be used for routing.

458 The details of this type of model will need to be worked out, but the potential for,
 459 and basic components and principles of, a DL-based integrated hydrology model are rel-
 460 atively clear. There is no fundamental limitation that precludes developing integrated
 461 DL hydrology models at multiple temporal and spatial scales. The questions that we an-
 462 ticipate are about what value will come from integrating different types of features and
 463 feature layers, and about how we might pre-train various feature layers to account for
 464 different types and scales of observational data in large, integrated models.

465 7.2 Theory-Informed Machine Learning

466 As mentioned in Section 5, there is a feeling among hydrologists and Earth scien-
 467 tists that models without explicit process representation might be unreliable under chang-
 468 ing conditions. Although we don't know if this is really true, one way to approach this
 469 is to integrate physical constraints or process-based theory into DL models. The goal is
 470 to extract as much information as possible from a combination of theory and data. This
 471 is not a new idea - Karpatne et al. (2017) called for *theory-guided data science*, which
 472 consists of efforts to integrate scientific consistency into generalizable models. Notably,
 473 members of that same group later collaborated on development of a DL model that is
 474 architecturally constrained to not violate prescribed monotonicity relationships (Daw et
 475 al., 2019).

476 A simple and general way to enforce conservation constraints (e.g., mass, energy,
 477 momentum) in a DL architecture is to L1-normalize a set of bounded ($\in [0, 1]$) activa-
 478 tion functions, and scale by the conserved quantity. This concept can be integrated into
 479 almost any type of neural network architecture, including into the long short term mem-
 480 ory networks used by Kratzert, Klotz, Herrnegger, et al. (2019) and Kratzert, Klotz, Shalev,
 481 et al. (2019). This concept is illustrated in Figure 7, and the result is a model that learns
 482 nonlinear input-state-output relationships that obey arbitrary and interacting conser-
 483 vation principles.

484 Another approach for directly combining process understanding with ML is to in-
 485 corporate the ML models inside of a dynamical systems model. A basic approach was
 486 outlined by Ghahramani and Roweis (1999), where - effectively - an empirical model is

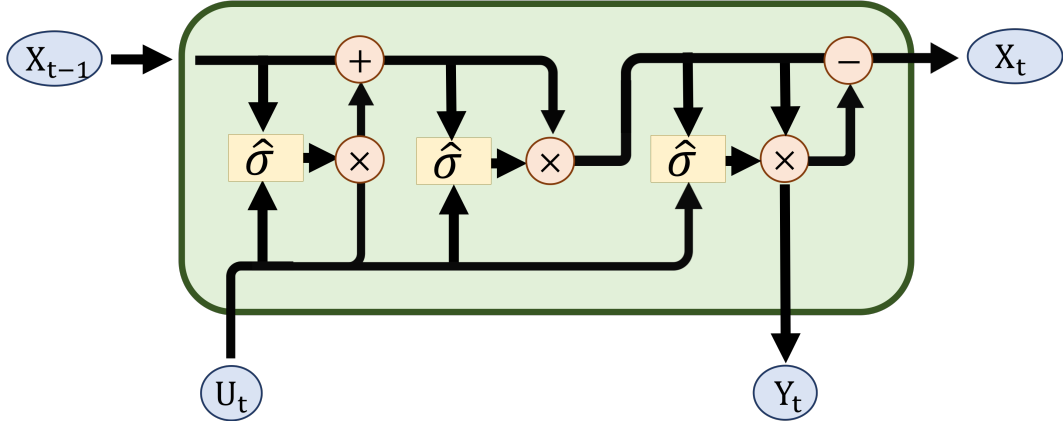


Figure 7. A time-recurrent deep learning network that is architecturally constrained to conserve mass, energy, and/or momentum. U_t are time-dependent inputs, Y_t are time-dependent outputs, and X_t is a vector of N memory states in the network, $\hat{\sigma}$ represents a set of N L1-normalized sigmoid activation functions that produce a set of real values in $[0,1]$ that sum to unity. These are scaled by the conserved quantities (in the inputs and states) so that the total sum of the time-history of inputs plus outputs is always equal to the total sum of the system state. There are three sets of ‘gates’ in this network - an input gate that moves mass (energy, momentum) from inputs to states, a reshuffling gate that moves mass (energy, momentum) between states during each individual timestep, and an output gate that moves mass (energy, momentum) from states to outputs at each timestep.

487 trained on the analysis states resulting from data assimilation (e.g., by a Kalman-type
488 filter). We can generalize this idea as follows:

Suppose that we have a dynamical systems model that solves a set of PDEs:

$$\frac{dX}{dt} = f(X, U, \theta), \quad (1)$$

where X are modeled system states, U are time-dependent boundary conditions, θ are model parameters, and function $f(\cdot)$ is the total divergence (inputs less outputs). A discrete-time approximate solution might then be:

$$X_t = f^*(X_{t-1}, U_t, \theta). \quad (2)$$

We can augment the $f^*(\cdot)$ state-transition function with a learned component, $g^*(\cdot)$, as:

$$X_t = f^*(X_{t-1}, U_t, \theta) + g^*(X_{t-1}, U_t, \theta). \quad (3)$$

489 where $g^*(\cdot)$ is any ML model. As above, $g^*(\cdot)$ can itself be probabilistic so that equa-
490 tion 3 is a discrete-time solution to a set of stochastic PDEs. The challenge is to learn
491 the $g^*(\cdot)$ function given that we can’t expect to have direct observation pairs (X_t, X_{t-1})
492 of all system states to use for supervised learning. As an example, Nearing and Gupta
493 (2015) applied the data assimilation approach by Ghahramani and Roweis (1999) to the
494 HyMod conceptual rainfall runoff model, and Pelissier et al. (2020) applied a similar tech-
495 nique to the Noah-MP land surface model for soil moisture accounting.

496 Another example of potential for theory-guided data science in hydrological work-
497 flows is for data assimilation itself. Significant information loss often results from assign-
498 ing the distributions and parameters of a probability-based assimilation algorithm (Near-
499 ing, Yatheendradas, et al., 2018) and many assimilation algorithms require that the model

500 and observation be in the same climatology (S. V. Kumar et al., 2012), meaning that
 501 these algorithms only treat stochastic error. One potential way to mitigate these prob-
 502 lems is to use ML to learn relationships between model states and assimilated observa-
 503 tions (e.g., Kolassa et al., 2018). As an example of this, Nearing (2013) derived the fixed-
 504 form of the Kalman-type gain and its associated adjoint that results from assimilating
 505 with a Gaussian process observation operator.

506 We see theory-guided data science, and more specifically, physics-informed ML, as
 507 a likely strategy for simultaneously leveraging what we do know from scientific theory
 508 about catchment behavior with the now-undeniable ability of DL for extracting patterns
 509 and information directly from data. There is some indication that this might be useful:
 510 Figure 8 shows a comparison between the performance of a DL model applied to CAMELS
 511 basins vs. a calibrated conceptual model. This data is from Kratzert, Klotz, Herrnegger,
 512 et al. (2019), and the takeaway message is that while the DL model is better over-
 513 all, it is not better everywhere. Kratzert, Klotz, Herrnegger, et al. (2019) could not find
 514 any relationship(s) between observable catchment characteristics and the difference in
 515 performance between these two models, but it is nevertheless apparent that there is at
 516 least the potential to improve by adding some elements of hydrologic theory to the DL
 517 architecture.

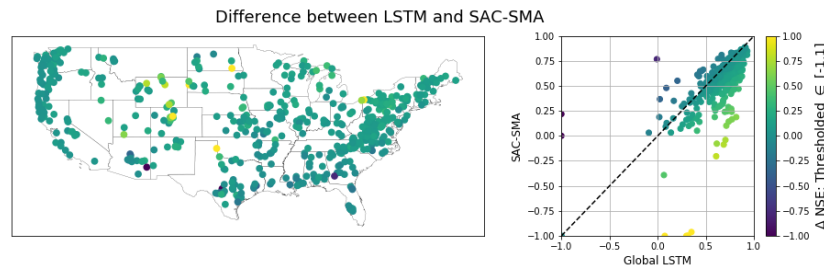


Figure 8. An illustration from Kratzert, Klotz, Herrnegger, et al. (2019) that compares a deep learning model (LSTM) against a calibrated conceptual model (SAC-SMA) over 531 CAMELS basins. The deep learning model does better on average, but not in every catchment, indicating that there is at least potential to improve by incorporating some of the conceptual constraints from SAC-SMA.

518 7.3 Skip the Hydrologist

519 Clark et al. (2016) gave an account of the sources of uncertainty (information loss)
 520 in a hydrological modeling chain. These are things like uncertainty in meteorological forc-
 521 ings from global circulation models (GCMs), downscaling forcings to the watershed scale,
 522 errors in the hydrological model structure, parameter uncertainty, etc. Each of these rep-
 523 represents a step in a chain of information from the GCM dynamical core (i.e., Navier-Stokes
 524 approximations and data assimilation) to streamflow or other hydrological variables. Ev-
 525 ery step in this modeling chain introduces uncertainty. DL has the potential to let us
 526 skip at least several steps in this type of modeling chain by developing relationships di-
 527 rectly between high-quality data sources.

528 Take as an example the largest source of hydrological error, which is typically precipi-
 529 tation data. This is true whether we are using the output of weather or climate mod-
 530 els, interpolated gauge data, or remote sensing data from radar and/or satellites. The
 531 problem is exacerbated by downscaling. The major precipitation-related uncertainty in
 532 a global circulation model is due to parameterization of sub-grid cloud formation pro-
 533 cesses. There have been recent successes using ML to parameterize cloud physics and

534 cloud formation (e.g., Gentine, Pritchard, Rasp, Reinaudi, & Yacalis, 2018), which could
 535 help mitigate these issues to some extent, but we still have to feed these uncertain pre-
 536 cipitation fields into a hydrology model that is subject to both parameter and structural
 537 uncertainties.

538 We could think about the problem in a different way. The four-dimensional pres-
 539 sure, wind, and temperature fields that result from Euler solutions in the dynamical cores
 540 of GCMs are relatively accurate, at least as compared with the accuracy of parameter-
 541 ized precipitation fields. We could, in principle, use DL to extract information directly
 542 from states of the dynamical core about terrestrial hydrological variables. For example,
 543 we could in principle develop four-dimensional convolutions to regress directly from GCM
 544 fields and digital elevation maps to pixel classifiers over satellite-derived maps of flood
 545 inundation, and thereby skip sources of information loss from (i) sub-grid convection pa-
 546 rameterizations, (ii) GCM downscaling, (ii) lack of scale-relevant theories of watersheds,
 547 (iii) parameter equifinality, (iv) rating curves, etc. It is possible (perhaps likely) that this
 548 type of model would give more accurate inundation forecasts at similar lead times rela-
 549 tive to state-of-the-art hydrology models, since this would skip uncertainties related to
 550 cloud physics parameterizations, downscaling, watershed parameterizations, etc. All of
 551 these things could be learned implicitly by a DL model.

552 The point is that DL offers at least the potential to make societally-relevant hy-
 553 drological forecasts without any type of hydrological model or hydrological process un-
 554 derstanding at all. Because DL allows for almost arbitrarily complex relationships, and
 555 has demonstrated to extrapolate well out-of-sample, it might be the case that success-
 556 ful water resources and water hazard predictions might not require anything that looks
 557 even like a simple hydrology model. This is all speculative, but the point is that the idea
 558 about hydrological understanding being necessary for reliable forecasting discussed in
 559 Section 5 may not be true even in the most superficial sense. This is an extreme and hy-
 560 pothetical example, but one that is worth (1) trying experimentally, and (2) being aware
 561 of as we calibrate our expectations about the role of hydrological theory and hydrolog-
 562 ical science in the context of big data and machine learning.

563 **7.4 Observations and Benchmarks**

564 Beven (2006b) proposed that the search for closure schemes at the watershed scale
 565 is the *second* most important problem in the discipline, with the most important being
 566 to improve observation capabilities. We agree completely. As was the case in 1987, the
 567 first and foremost job of hydrologists are and will continue to be related to improving
 568 observational capacity. The approaches discussed in this article only increase the need
 569 for observation data related to as many aspects of the water cycle as possible.

570 Shen et al. (2018) noted that past progress in the field of machine learning can be
 571 partially attributed to the culture of using public data sets and benchmarking new meth-
 572 ods against previous state-of-the-art. There have been calls for consistent practices re-
 573 lated to hypothesis testing, model intercomparison, and model rejection (e.g., Beven, 2018).
 574 While some of the philosophical counter-arguments to this are compelling (e.g., Baker,
 575 2017; Nearing et al., 2020; Nearing, Tian, et al., 2016), without *some* community stan-
 576 dard for benchmarking it is difficult to track progress in the field in an objective way.

577 This means that we need two things. First are better centralized data repositories.
 578 The community is aware of this (H. Gupta et al., 2014; Shen et al., 2018) and there are
 579 several such efforts happening in the field right now (e.g., Addor et al., 2017; Hoffman,
 580 Riley, Randerson, Keppel-Aleks, & Lawrence, 2016; Newman et al., 2015). We expect
 581 that this issue will sort itself out in the near future. Still, our opinion is that one of the
 582 best investments that could be made in the discipline right now is to develop standard-
 583 ized and easily accessible big data repositories. The second thing we need is the willing-
 584 ness to use those data repositories. Just like in previous decades when the community

585 responded to calls for making uncertainty quantification required for every modeling study
 586 (Pappenberger & Beven, 2006), we need a community standard that requires all new mod-
 587 eling papers to include large-scale benchmarking against standard, centralized data sets.

588 Hydrological modeling is currently a field of ivory towers where legacy and affil-
 589 iation guide the choice of model Addor and Melsen (2019) as opposed to empirical rigor
 590 Beven (2018). Different modeling groups largely work on their own models, and while
 591 there have been ad hoc intercomparisons (e.g., Best et al., 2015; van den Hurk et al., 2011),
 592 this is not routine and the hydrology community does not keep a list of current perfor-
 593 mance scores on standard test problems, as is standard in other communities (e.g., CMIP,
 594 ML-Perf, etc.).

595 8 A White Whale

596 During the community contribution phase of the IAHS ‘Unsolved Problems in Hy-
 597 drology’ effort (Blöschl et al., 2019), one of the suggested questions was: “*Does Machine*
 598 *Learning have a real role in hydrological modeling?*” In contrast, we suggest that the ex-
 599 istential question for our discipline right now is: “*What role will hydrological science play*
 600 *in the age of machine learning?*” van den Hurk et al. (2011) challenged that “*it must*
 601 *be demonstrated that the model physics actually adds information to the prediction sys-*
 602 *tem.*” This is exactly the question that needs to be answered in order to understand how
 603 and where hydrological theory has a role to play in a world dominated by data. We see
 604 at least potential for deep learning to help address this by allowing us to decouple dif-
 605 ferent parts of hydrological theory while still retaining scale-relevant predictive systems
 606 learned (partially) from data.

607 Very likely, the future of hydrology will be a mix of AI and physics-based approaches,
 608 but we have a hard time envisioning a future where transformative data science approaches
 609 like DL become simply another tool in the hydrologist’s toolbox. We see it as much more
 610 likely that hydrological domain knowledge will become an integral part of guiding and
 611 developing fundamentally AI-based systems and analyses (e.g., Section 7.2).

612 Hydrology has roots - at least in part - as a branch of civil engineering. Klemeš (1986a)
 613 argued that “*practices of bad science in hydrology cannot be blamed on engineers and other*
 614 *decision makers who ‘need numbers.’ For if these numbers are not to be based on sound*
 615 *hydrologic science but only on manipulations of arbitrary assumptions and concepts, hy-*
 616 *drologists are not needed. Engineers can do such a job much better themselves since they*
 617 *at least can tailor the assumptions to the particular projects and, not mistaking them for*
 618 *scientific truth, will treat them accordingly in the decision process.*” The situation has
 619 not changed much in the 34 years since this was written: our ability to extract numbers
 620 (predictions) from data is advancing rapidly, but we have not improved very much our
 621 ability to make predictions from anything resembling hydrologic theory. While our mod-
 622 els become increasingly complex, a well-calibrated Sacramento model is still one of the
 623 best in discipline.

624 The reason that we think this is an *existential* challenge is because we see hydro-
 625 logical science becoming increasingly decoupled from state-of-the-art hydrological infor-
 626 mation systems. Major development groups at governmental institutions internation-
 627 ally continue to dedicate the large majority of effort to the traditional models that have
 628 never benchmarked well against ML (e.g., Abramowitz, 2005; Best et al., 2015; Kratzert,
 629 Klotz, Herrnegger, et al., 2019; Nearing, Ruddell, et al., 2018). As far as we can tell, these
 630 models are dead on arrival. Barring some major fundamental theoretical discovery or
 631 innovation, there is essentially no chance that any incremental advancements will allow
 632 these models to catch up to the state-of-the-art hydrological predictions. Simultaneously,
 633 there has not been any serious or systematic investment into AI-based hydrology at a
 634 meaningful scale, and from what we can see (e.g., see Section 5) there is still strong re-

sistance in the hydrology community toward adopting these approaches in a serious and fundamental way. Coupled with the fact that DL experiments demonstrate that hydrologists lack even a basic understanding of why their models fail (Section 2), this causes us to worry.

Our fear is that if the hydrological sciences community refuses to make a serious investment into the technology that works, then someone else will. This will mean a further decoupling between hydrological science (such as it is) and the societal value that this science is supposed to support. To be clear, the current authors do *not* want to see that happen, but we are not impressed with the reaction we are seeing in the community. Our message in this opinion piece is to stop assuming that the world needs our theories and expertise, and start demonstrating - quantitatively and systematically - the value of individual components of that expertise against the backdrop of a growing importance of big data.

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