

# Holocene deglaciation and glacier readvances on the Fildes Peninsula and King George Island (Isla 25 de Mayo), South Shetland Islands, NW Antarctic Peninsula

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## **Abstract**

To provide insights into glacier-climate dynamics of the South Shetland Islands (SSI), NW Antarctic Peninsula, we present a new deglaciation and readvance model for the Bellingshausen Ice Cap (BIC), Fildes Peninsula, King George Island/Isla 25 de Mayo (KGI) ~62°S. Deglaciation on KGI began after c. 15 ka and had progressed to within present-day limits on the Fildes Peninsula, its largest ice-free peninsula, by c. 6.6–5.3 ka cal BP. Probability density phase analysis of chronological data constraining Holocene glacier advances on KGI revealed up to eight 95% probability ‘gaps’ during which readvances could have occurred. These are grouped into four stages – Stage 1: a readvance and marine transgression, well-constrained by field data, between c. 7.4–6.6 ka cal BP; Stage 2: four probability ‘gaps’, less well-constrained by field data, between c. 5.3–2.2 ka cal BP; Stage 3: a well-constrained readvance between c. 1.7–1.5 ka; Stage 4: two further readvances, one less well-constrained by field data between c. 1.3–0.7 ka cal BP (68% probability), and a ‘final’ well-constrained readvance after <0.7 ka cal BP. The Stage 1 readvance occurred as colder and more negative Southern Annular Mode (SAM)-like conditions developed, and marginally stronger/poleward shifted westerly winds led to more storms and precipitation on the SSI. Readvances were more frequent after c. 5.3 ka cal BP, driven by reducing spring/summer insolation at 62°S and negative SAM-like conditions, but weaker (equatorward shifted) westerlies over the SSI led to reduced storminess, restricting readvances close to present day limits. Late Holocene readvances are anti-phased with sub-aquatic moss layers in lake records unaffected by glaciofluvial inputs. Retreat from ‘Neoglacial’ glacier limits and the recolonisation of lakes by sub-aquatic moss in the post-bomb era (>1950 CE) is associated with recent warming/more positive SAM-like conditions.

## **Keywords**

Palaeoenvironmental change, chronology, geomorphology, diatoms, tephra, Southern Hemisphere Westerly winds

## Introduction

Field-based reconstructions of Antarctic ice sheet and ice cap deglaciation after the Last Glacial Maximum (c. 26–18 ka) and into the Holocene are important components of ice sheet models used to examine the impact of future ice-melt on rates of regional-global sea-level change (Sugden and Clapperton, 1977; Ó Cofaigh et al., 2014; Lindow et al., 2014; Simms et al., 2018, 2019; Kaplan et al., 2020; Nývlt et al., 2020; Oliva et al., 2020). Despite numerous studies, the early-mid Holocene deglaciation and the number, timing, and extent of mid-late Holocene glacier readvances on the northern Antarctic Peninsula (NAP) and South Shetland Islands (SSI) remains debated (Hall, 2007; Hall, 2009; Simms et al., 2011; Watcham et al., 2011; Simms et al., 2012; Ó Cofaigh et al., 2014; Emslie et al., 2019; Oliva et al., 2019).

Ice-marginal and ice-free coastal areas on the NAP and SSI contain an abundance of readily accessible morphostratigraphic features (e.g., raised beaches), lake sediment records, and glacially transported erratics which can be used to constrain past glacier dynamics. The glacial history of the Fildes Peninsula in the south-west of King George Island (KGI)/Isla 25 de Mayo, the largest of the South Shetland Islands (Figure 1), has been determined by cosmogenic isotope exposure dating of bedrock, radiocarbon and optically stimulated luminescence (OSL) dating of organic remains embedded in raised beaches and moraines of the c. 1250 km<sup>2</sup> Bellingshausen Ice Cap (BIC; also known as the Collins Ice Cap), as well as radiocarbon ages from lake sediments and terrestrial stratigraphic sequences overlying coarse grained sediments and diamictons (Figure 1) (Sugden and John, 1973; Pallàs et al., 1997; Bentley et al., 2005; Seong et al., 2008; Hall, 2010; Watcham et al., 2011; Simms et al., 2012; 2021).

[insert Figure 1.]

Isolation basins, raised beaches, and marine abrasion terraces used to constrain glacio-isostatic adjustment (GIA) (uplift) and changes in relative sea level (RSL) are often regarded as the most reliable datasets for establishing deglacial unloading and reloading associated with glacier readvance(s) (Bentley et al., 2005; Johnson et al., 2022). The SSI have a complex geological structure due to their position above a former subduction zone and adjacent to an extending basin in Bransfield Strait (Christeson et al., 2003) and variations in elastic strength and underlying viscosity have not yet been captured in GIA models (van der Wal et al., 2015). Consequently, RSL data from the SSI are more complex than other locations across Antarctica,

and several different RSL curves have been suggested (Pallàs et al., 1997; Bentley et al., 2005; Hall, 2010; Roberts et al., 2011; Watcham et al., 2011; Simms et al., 2012; 2021; Johnson et al., 2021).

While some GIA models (e.g., Whitehouse et al., 2012) infer changes in ice loading that are consistent with RSL reconstructions based on well-dated isolation basin records from the Fildes Peninsula (Watcham et al., 2011; Roberts et al., 2017), the coarse spatial resolution of GIA models means they are not currently able to capture smaller scale isostatic responses and RSL reversals implied by field data (Simms et al., 2012; Whitehouse et al., 2012; Simms et al., 2018; 2021; Johnson et al., 2021). Spatially variable responses of glacio-isostatic uplift are thought to have occurred across the SSI and AP during the mid-late Holocene (Fretwell et al., 2010; Zurbuchen and Simms, 2019), necessitating new data collection and new methods of constraining glacier readvances.

In this study, we examined the deglaciation and readvance history of the BIC outlet on the Fildes Peninsula using a combined geomorphologic, stratigraphic, chronological, palaeolimnological and statistical approach to provide new constraints on glacier readvance(s) during the mid-late Holocene. New and published data were used to test the hypothesis that Holocene retreat of the BIC was discontinuous and interrupted by more than one glacial readvance during the mid-late Holocene, rather than a continuous process of glacial retreat.

To achieve this, first, we obtained new minimum Beryllium-10 ( $^{10}\text{Be}$ ) cosmogenic isotope exposure ages for deglaciation from granitic erratics on the NW marine platform of the Fildes Peninsula and the BIC foreland (Figure 1). Second, we obtained new radiocarbon ages from terrestrial and marine macrofossils that were living on the BIC foreland on the Fildes Peninsula before glacier readvances embedded them into moraines, which represent maximum ages of glacier advance. Third, we undertook multi-proxy analyses on a new lake sediment core extracted from Kiteschee Lake and compared these results to basal ages and multiproxy data from previously published lake records on the Fildes Peninsula to determine when deglaciation occurred in different sectors of (inner, mid, outer; Figure 1) and how Holocene glacier-climate dynamics impacted on lake ecosystems.

With the addition of these new datasets, the amount of chronological data constraining Holocene glacier fluctuations on KGI were large enough to use probability density phase



analysis to identify ‘probability gaps’ when glacier readvances could have occurred. In summary, we identified eight occasions in the mid-late Holocene when glacier readvances could have occurred and examined mechanisms driving deglaciation and glacier readvance(s) on the Fildes Peninsula and across KGI.

## **Regional setting & Study area**

### **Climate setting**

The SSI archipelago is a ~230 km long, ~35 km wide active volcanic arc of 11 major islands, with the Fildes Peninsula located at ~62° S, 58–59° W and ~130 km northwest of the Antarctic Peninsula (Figure 1B). The SSI are located south of the Polar Front at 50–55°S, where the southern limb of the Southern Hemisphere westerly winds (SHW) belt converges with the polar cell (Figure 1A). Sandwiched between the high pressure subtropic and Antarctic anticyclones, the SSI island chain has a milder and more humid (wetter) maritime climate than the Antarctic continent with more than ~1000 mm yr<sup>-1</sup> of precipitation, which maintained a stable to slightly negative glacier mass balance for the BIC and an equilibrium altitude (ELA) of ~ 150 m a.s.l. between 1970–90 CE (Turner et al., 2002; Bentley et al., 2009; Falk et al., 2018). Intense cyclonic activity through the Drake Passage, associated with the SHW to the north, and the low-pressure circumpolar belt to the south (~66°S) means the SSI experience cloud cover for ~80% of the year (King and Turner, 1997; Bañón et al., 2013).

The annual mean air temperature on KGI between 1948 and 2011 CE was -2.5 °C (Kejna et al., 2013). Mean annual, summer and winter air temperatures at the Russian Bellingshausen Station, on the Fildes Peninsula, between 1968–2015 CE (Olivia et al., 2017) and (1968–2021 CE) (BAS READER project – Reference Antarctic Data for Environmental Research <https://www.bas.ac.uk/project/reader/>; Turner et al., 2004) were -2.3°C (-2.3±3.1°C), 1.1°C (1.1±0.8°C) and -6.1°C (-5.9±1.7°C), respectively, with mean summer air temperatures above zero and snow free conditions (at sea level) up to four months between December and March (Michel et al., 2014). Mean annual, summer and winter wind velocities recorded at Bellingshausen (1968–2021 CE) were 26.5±3.5 km h<sup>-1</sup>, 23.8±2.2 km h<sup>-1</sup>, 27.3±3.2 km h<sup>-1</sup>, respectively. Storms and ~100 km h<sup>-1</sup> winds are common on the SSI, and wind directions are predominantly SW/W/NW to NE/E/SE (34% W/NW to E/SE), but rarely NE-SW axis (BAS READER project).

A warming trend across the AP and SSI between the 1950 and 1999 CE drove increased rates of glacier retreat, ice-shelf collapse (Vaughan et al., 2003; Meredith and King, 2005) and elevated the ELA on the SSI to above 200 m above present sea level (henceforth, m a.s.l.) (Falk et al. 2018). Similar to most other stations on the AP and SSI, temperature data from the Bellingshausen Station exhibit decadal scale variability, which has been linked to the El Niño Southern Oscillation (ENSO) (Oliva et al., 2017). Warming initiated ~5 years prior to El Niño in 1982–3 CE and 1997–8 CE was followed by a cooling trend to 2015 CE and a return to a warming trend between 2015–21 CE (Figure 1E). Although a shift to a cooling trend was observed across the northern AP (NAP) between 1999–2015 CE, record summer temperatures have since been measured on the northeastern AP and attributed to an enhanced foehn effect (Turner et al., 2016; Oliva et al., 2017).

### **Geology and Glaciation**

King George Island (KGI), the largest of the South Shetland Islands, has been subdivided into three tectonic regions: the Fildes Block (Fildes Peninsula), Barton Horst (Barton and Weaver Peninsulas) and the Warszawa Block (Potter Peninsula) (Smellie et al., 1984; Watcham et al., 2011). Covering an area of 38 km<sup>2</sup>, the Fildes Peninsula is the largest ice-free area on the South Shetland Islands. It is composed of predominantly andesitic and basaltic lava bedrock, with some interbedded terrestrial sediments, including shales and conglomerates (Smellie et al., 1984). Most of the Bellingshausen Ice Cap (BIC) foreland was glaciated during the Last Glacial Maximum (pre-LGM) and consists of raised marine platforms up to 180 m a.s.l. of pre-LGM age and formed by basic volcanic rocks of Late Cretaceous to Paleogene-age (John and Sugden, 1971). The eastern coastline of the Fildes Peninsula is indented with coves, while the western coast is characterized by sheer sea cliffs and stacks (Hall, 2003).

The elevated island interiors of the SSI are characterized by numerous ice caps and permanent snowfields and the coastal fringes are some of the most extensive ice-free and permafrost areas in Antarctica (Braun et al., 2004; Ó Cofaigh et al., 2014) (Figure 1B). Chlorine-36 (<sup>36</sup>Cl) exposure ages from glacially striated bedrock on nearby Barton Peninsula from above c. 50 m a.s.l. show that progressive ice thinning on KGI, SSI started c. 15,000 years ago (Seong et al., 2008). On the northern half of the Fildes Peninsula, the BIC is bounded by prominent moraines, referred to as Shetland I (Hall, 2007). A prominent, and largely featureless, marine platform of pre-Holocene age at c. 40–35 m a.s.l. on the north-western Fildes Peninsula, is more than 7 km long and 1 km wide (John and Sugden, 1971) (Figure 1). This feature is flat to gently

undulating, seaward sloping with a limited number of undated granite erratics on its surface (Barsch and Mäusbacher, 1986; Hall, 2003). As far as we are aware, granite erratics on the South Shetland Islands and BIC foreland have not been previously dated using the cosmogenic exposure dating method.

Radiocarbon ages obtained from marine-terrestrial transition sediments and raised beaches have been used to reconstruct past changes in local RSL at several sites across the SSI. These indicate that the thickest part of the SSI ice cap was probably centred on the now ice-free area of the Fildes Peninsula during the LGM (Watcham et al., 2011). The SSI ice cap was possibly separated from the AP ice sheet by the Bransfield Marginal Basin during the LGM, but more likely became independent during the Lateglacial–Interglacial transition (c. 15–12 ka) (John and Sugden, 1971; Fretwell et al., 2010; Watcham et al., 2011; Ó Cofaigh et al., 2014).

Basal lacustrine sedimentary sequences in the mid-southern part of the Fildes Peninsula have been dated to c. 10 ka cal BP (e.g., Jurasee Lake (Mäusbacher et al., 1989)) and were formed in over-deepened glacial basins. Several former marine embayments with basal sediment ages of up to c. 12 ka cal BP (Watcham et al., 2011) were transformed into freshwater (isolation) basins when the rate of isostatic uplift outpaced the declining rate of sea level rise during the early Holocene (Watcham et al., 2011). A well-defined Holocene marine high-stand of 18–15 m a.s.l. occurred on the Fildes Peninsula between c. 8 and 7 ka (Mäusbacher et al., 1989; Watcham et al., 2011).

## **Methods**

### **Cosmogenic nuclide surface exposure dating of erratics**

The cosmogenic isotope concentration accumulated in a superficial rock sample is proportional to the length of time exposed to cosmic rays at the Earth's surface (Balco, 2011). The cosmogenic exposure age produced depends on the production rate (i.e., the concentration of cosmogenic isotopes produced each year per gram of the relevant mineral), which varies spatially and temporally due to variations in atmospheric depth and geomagnetic field effects (Lifton, 2016). Cosmogenic  $^{10}\text{Be}$  has the most well-constrained production rate and can be measured at low concentrations (Balco, 2011), but glacially striated and quartz-rich, granitic bedrock and erratics required for  $^{10}\text{Be}$  analysis are rare in the basaltic-andesitic volcanic-arc environment of the South Shetland Islands.

All of the limited number of large boulders >50 cm in diameter deposited on the NW BIC glacial foreland as glaciers retreated from offshore LGM limits across were surveyed and classified. Differential GPS (dGPS; WGS84 ellipsoid) data for erratic boulders was obtained using a GPS Trimble Pathfinder ProXH. As geodetic reference, we used the landmark DALL 66019M002 (S62°14'16.335", W58°39'52.364", ellipsoidal height 39.376 m) at the Argentine Carlini base, c. 17 km from the sampled erratics. Post-processed uncertainties for the samples were less than 0.1 m in altitude and in the horizontal.

The only three large granitic boulders not incorporated into solifluction planes and till sheets of the BIC foreland were considered suitable for cosmogenic nuclide exposure dating (Table 1). The sampled erratics showed no signs of significant erosion, and their size meant it is very unlikely they had been overturned. Post-depositional movement was minimized by sampling erratics away from slopes or cliffs. Samples were taken with a hammer and chisel, removing the upper few centimetres of sub-horizontally exposed surfaces (Figure 2C) and the angle to the skyline was measured to calculate topographic shielding (Balco et al., 2008). Surface shielding due to snow cover was minimised by sampling from wind-exposed localities (cf. Johnson et al., 2012, 2017, 2020; Glasser et al., 2014; Lindow et al., 2014). Laboratory analysis for <sup>10</sup>Be cosmogenic nuclide surface exposure dating followed the procedures of Kohl and Nishiizumi, (1992) and Binnie et al., (2015) (Table 1, Supplementary Materials for details).

### **Radiocarbon dating of moraines and stratigraphic sections**

Moraines adjacent to the BIC were mapped and interpreted from field observations and satellite images (DigitalGlobe, Catalogue ID: 1030010020C0C900; Google Earth, 2006 and 2011). Contour lines are derived from the Antarctic Digital Database with elevation data ( $\pm 5$  m) obtained from the Instituto Antártico Uruguayo (1997). Radiocarbon ages from moraines and stratigraphic sections were obtained by Accelerated Mass Spectrometry (AMS) dating of marine mollusc shells, terrestrial mosses and seaweed layers embedded in sediments and represent maximum ages for BIC glacier readvance (Table 2). To reduce the risk of contamination with modern material, all samples for radiocarbon dating were taken from fresh and previously cleaned outcrops or sediment cores, packed in zip-lock plastic bags, and stored at 4°C. Sample preparation prior to radiocarbon dating analysis was undertaken at the Alfred Wegener Institute Helmholtz Centre for Polar and Marine Research and the British Antarctic Survey. Sample measurements and corrections for <sup>13</sup>C/<sup>12</sup>C ratio and calculation of

Conventional Radiocarbon Ages were undertaken by ETH Zürich and SUERC, Scotland (Table 2, Supplementary Materials for details).

Calibration of radiocarbon ages from marine samples was undertaken in Oxcal v. 4.4 using the Marine20 calibration curve (Bronk Ramsey, 2009; Heaton et al., 2020), and a newly recalculated local marine reservoir age offset ( $\Delta R$ ) of  $666 \pm 76$   $^{14}\text{C}$  years (Heaton et al., 2020), which represents the weighted mean  $\Delta R$  of four radiocarbon-dated marine samples collected prior to 1950 CE from the northern Antarctic Peninsula and Signy Island in the CALIB online Marine20 database (<http://calib.org/marine/>) (Table 2, S3; Supplementary Materials for details). Terrestrial samples, Kiteschsee Lake samples and previously published data from Yanou Lake (Roberts et al., 2017), were (re)calibrated using the Southern Hemisphere SHCal20 calibration curve (Hogg et al., 2020) in Oxcal v. 4.4 (Bronk Ramsey, 2009). There were no significant statistical differences (i.e., beyond two-sigma error ranges) between Holocene radiocarbon ages calibrated using the SHCal13/Marine13 (Hogg et al., 2013) and the ShCal20/Marine20 curves (ShCal13 vs 20:  $n=19$ , mean difference  $\pm 1\sigma = -4 \pm 8$  years; Marine:  $n=8$ , mean difference  $\pm 1\sigma = -45 \pm 8$  years); hence, comparisons with published data and age models calibrated using the SHCal13 and Marine13 curves remain valid.

### **Kiteschsee Lake record**

Kiteschsee lake ( $62^\circ 11'36.55''\text{S}$ ,  $58^\circ 57'59.93''\text{W}$ ) is located at c. 15 m a.s.l. in the eastern-central region of the Fildes Peninsula, c. 550 m from the east coast and approximately 2.6 km from the margin of the BIC (Figure 1D, 2B). It is adjacent to the Davies Heights ( $>100$  m a.s.l.), a previously glaciated part of the former raised marine platform in the mid-central area of the Fildes Peninsula, which supplies meltwater directly into the eastern end of the lake (Figure 1D). Marine sediments at the base of a previously studied sedimentary record from Kiteschsee Lake (Mäusbacher et al., 1989) show that much of the low-altitude central area of the Fildes Peninsula remained below sea-level until c. 7–6 ka (Mäusbacher et al., 1989).

*Lithology, geochemistry and chronology:* We undertook multi-proxy analyses (diatom, grain size, geochemical and sedimentological analysis) on a 77 cm-long sediment record extracted from the flat-bottomed eastern basin depocentre of Kiteschsee Lake (Figure 4, S1–3) and compared data obtained with published lake records from the Fildes Peninsula. High resolution X-ray fluorescence (ITRAX) core scanning (XRF-CS) of bulk sediment was performed at

contiguous 0.02 cm (200  $\mu\text{m}$ ) intervals following procedures in Davies et al. (2015) and Roberts et al. (2022) (Supplementary Materials for details).

A chronology for the Kiteschsee Lake sediment record was based on two aquatic moss ages, six bulk sediment AMS radiocarbon ages, and four tephra correlation ages and established using a Bayesian age-depth modelling in BACON v. 2.5 in R (Blaauw and Christen, 2011) (Figure 4, Table 2). Volcanic glass-shard counts, and electron probe microanalysis (EPMA) data and tephrastatigraphic correlation with the nearby Yanou Lake record were used to improve the final age-depth model (Blockley et al., 2005; Roberts et al., 2017) (Figure 4, Supplementary Material for details). Reporting of radiocarbon data and age-depth modelling parameters follows recommendations in Lacourse and Gajewski (2020). Calibrated ages in the text have been rounded to the nearest hundred years in the text to reflect realistic total (internal and external) uncertainties in radiocarbon dating.

*Diatom analysis:* Since the early Holocene isolation history of Kiteschsee Lake is well-established (Mäusbacher et al., 1989), and diatom smear-slide screening tests revealed that samples from 77–47 cm depth had very low diatom concentrations, quantitative analysis focussed on reconstructing high-resolution diatom compositional changes in the uppermost 47 cm (late Holocene) where diatom preservation was sufficient for meaningful environmental interpretation (Figure 4).

Quantitative diatom counting and analysis was undertaken at 1–2 cm intervals on sediments between 47 cm and 5 cm depth (3,680 – -63 a cal BP) and related to XRF-CS data using Principal Component Analysis (PCA) (Figure S10, Supplementary Materials for details). Diatom compositional turnover ( $\beta$ -diversity in Figure 4, 5) was estimated down-core using Detrended Canonical Correspondence Analysis (DCCA) for the most abundant diatoms (>2%). To infer changes in chlorophyll-a shown in Figure 4, we applied the diatom-chlorophyll-a training set developed for the Antarctic Peninsula using data from 61 lakes (Jones and Juggins, 1995). Diatom-based transfer functions were applied in C2 (Juggins, 2007) using simple weighted averaging (WA) and weighted partial least squares (WA-PLS) algorithms. The WA-Inverse transfer function was chosen to reconstruct chlorophyll-a as the RMSE, and average bias were low and displayed a strong relationship ( $R^2$ ) between measured and diatom-inferred chlorophyll-a (Supplementary Materials for details).

Constrained cluster analysis (CONISS) with broken stick analysis was applied to square root transformed datasets using R packages Vegan and Rioja to define a significant number of lithofacies units and diatom zones and their boundary positions (Juggins, 2007; Juggins, 2012; Oksanen, 2014).

### **Data compilations and analysis**

New and published chronological data constraining Holocene glacier readvances from the Fildes Peninsula (n=43), Potter Peninsula (n=29), and KGI (n=76) were compiled and a non-parametric phase model (i.e., a probabilistic version of the Oxcal SUM command) was applied to each dataset using the Bchron v. 4.7.6 in R (Haslett and Parnell, 2008; Parnell, 2021). This analysis provides a Gaussian mixture prior distribution density phase age-range estimate, at 0.68 and 0.95 probability level, constraining periods when glacier readvances could have occurred on the Fildes, Potter and Barton Peninsulas on KGI (FP, PP and BP in Figure 3; Table 3). Density phases produced by similar analysis of published SSI-wide data were less well-defined, likely reflecting local glacier influences and larger errors for some data, and not considered further. The Rcarbon package v. 1.4.2 (Bevan, 2021) was also used for data analysis and plotting. Mapping was undertaken in ARC-GIS and data was analysed and plotted using the Tidyverse suite of packages in R v. 4.1.0/Rstudio v. 1.4.1717, Sigmaplot v. 14.0, C2 (Juggins, 2007), and MATLAB v. R2021a.

## **Results**

### **Cosmogenic nuclide surface exposure dating of erratics**

The three erratic granitic boulders on the NW marine platform shown in Figure 3 gave an error-weighted mean  $\pm$  internal (external) error cosmogenic surface nuclide exposure age of  $6,780 \pm 220$  (590) years (chi-squared, p-value 0.9718), based on the nearby Patagonian production rate (Kp in Table 1) calibration dataset, and  $6,580 \pm 209$  (442) years (chi-squared p-value 0.9715) based on the global calibration dataset (Ww in Table 1) (Borchers et al., 2016, both using the “LSDn” scaling scheme (Table 1). The relatively small standard deviation of the three ages of  $\pm 60$  years is a favourable indicator of low analytical uncertainties but omits any systematic bias in the age related to, for example, production rate. Henceforth, we base our discussions on the LSDn scaled age estimates ( $\pm$ external errors) and their weighted mean age of  $6,580 \pm 440$  years (Figure 2; Table 1), as this provides a minimum age constraint for ice retreat coupled with a realistic estimate of external uncertainty, taking production rate

uncertainty into account. However, the close agreement of the three exposure ages and the overlap between ages produced using the three different production rate calibration datasets shown in Table 1 implies the samples had negligible pre-exposure and that they were exposed simultaneously on the NW foreland during a period of (potentially) rapid mid-Holocene glacier retreat.

[insert Table 1.]

[insert Figure 2.]

### **Radiocarbon dating of moraines and stratigraphic sections**

Twenty-five new radiocarbon ages were obtained from ten samples in the Artigas Beach–Valle Norte sections of the Shetland I moraine and its associated glaciofluvial and marine sediments (Figure 2B, 3A–C; Table 2). Shells embedded in marine sandy to silty sediments reworked into the central Valle Norte sector of the Shetland I moraine (Figure 2B) returned near infinite radiocarbon ages >40 ka cal BP, implying significant reworking of older glaciomarine deposits by the BIC (Table 2). In contrast, shells reworked into till in the Artigas Beach southern sector (Figure 2B) have a mean calibrated age range of 6.6–5.4 ka cal BP (Table 2), with two distinct 0.95 probability density age phases of 6.8–6.5 and 6.3–5.3 ka cal BP (Figure 3C; Table 3). The oldest of these ages are broadly coeval with cosmogenic exposure ages of the granite erratics (Figure 3C), which represent a minimum retreat age from the northern and central sectors of the inner BIC foreland, while the youngest are consistent with basal (deglaciation) ages from the inner Fildes Peninsula lake sediment records in the Valle Norte–Artigas Beach sectors (Figure 3C, E) (Watcham et al., 2011; Roberts et al., 2017).

Terrestrial moss fragments in glaciofluvial and glacial sediments from the Shetland I moraine dated in this study have mean age ranges of 2.0–1.2 ka cal BP into two well-defined 95% probability density phases of 2.1–1.7 ka cal BP and 1.4–1.1 ka cal BP (n=13) (Figure 2B, 3A, D; Table 3), suggesting emplacement by limited late Holocene glacier readvances between 1.7–1.4 and/or after 1.1 ka cal BP (Figure 3A).

Non-parametric 95% probability density phase modelling of all maximum age constraints from terrestrial mosses and shells emplaced in Shetland I moraines from this study (n=25) and that of Hall (2007) (n=43 in total) revealed up to six ‘probability gaps’ when the readvance of the



Shetland I moraine could have occurred: 7.3–6.8 ka cal BP, 5.3–3.5 ka cal BP, 3.1–3.0 ka cal BP, 2.6–2.2 ka cal BP, 1.7–1.4 ka cal BP, <0.2 ka cal BP (Figure 3A–C; Table S3). Analysis of new and existing chronological constraints on glacier advances from the Fildes Peninsula (Hall, 2007; this study), Potter Peninsula (Heredia Barión et al., submitted) and Barton Peninsula (Oliva et al., 2019) (n=80) identified seven 95% ‘probability gaps’ during which readvances could have occurred (7.4–6.6 ka cal BP, 5.3–4.8 ka cal BP, 4.5–3.9 ka cal BP, 3.3–3.0 ka cal BP, 2.6–2.5 ka cal BP, 1.7–1.5 ka cal BP, and <0.2 ka cal BP) and one 68% ‘probability gap’ (1.3–0.7 ka cal BP) (Figure 3D, E; Table S3). Aquatic moss with post-bomb <sup>14</sup>C ages is a common feature at sediment-water interfaces in lakes on Fildes Peninsula lakes (Figure 1) and likely reflects warmer conditions on KGI since the mid-C20th. Late Holocene ‘readvance probability gaps’ are broadly anti-phased with maximum to minimum calibrated <sup>14</sup>C age ranges constraining aquatic moss deposition in lake records from Fildes Peninsula (red bars in Figure 3E).

[insert Table 2.]

[insert Figure 3.]

### **Kiteschsee Lake record**

The new Kiteschsee Lake record shown in Figure 4A, B provides a continuous and well-defined mid–late Holocene record of palaeoenvironmental change and volcanic activity. It is largely devoid of soliflucted ash and catchment reworking that disrupts sedimentation in other lakes on the Fildes Peninsula and KGI, e.g., Yanou Lake (Figure 4C) (Watcham et al., 2011; Roberts et al., 2017).

*Lithology, geochemistry and chronology:* Downcore changes in XRF-CS Ti, K, Ca, and sand content reflect catchment erosion of (volcanic) bedrock and tephra deposition. These and tephra glass-shard counting revealed the position of several prominent visible volcanic ash (tephra) layers (T1–T7 in Figure 4, S2, S3, S4). Incoherence/coherence scatter ratios (inc./coh.) are proxy for organic content and were also used to define the precise positions of the visible tephra layers in Figure 4, S2, S3. The two most prominent tephra layers at 33 cm (mean; 95% range: 1,730; 2,420–1,470 a cal BP) (labelled T2 in Figure 4A–B) and 59 cm (5,600; 5,740–5,450 a cal BP) (T5 in Figure 4A–B) have a similar airfall thickness to the airfall components of the two most prominent mid–late Holocene tephra deposits in nearby Fildes Peninsula lake

records, e.g., Yanou Lake (Figure 4C, D) (Watcham et al., 2011; Roberts et al., 2017). Both layers contained rhyolitic glass whose shard geochemistry was similar to tephra deposits on the AP that have been linked to mid Holocene eruptions from the Deception Island volcano rather than South America (see Supplementary Materials for details).

Lithofacies units were defined by cluster analysis of XRF-CS data and are summarised as follows: Units 1–3, 77–60 cm (7.5–5.6 ka cal BP, mean modelled ages) are composed of poorly sorted olive-brown silt with three prominent fine black ash layers in Unit 2. After deposition of the T5 tephra (Unit 4: 60–58 cm; 5.6–5.5 ka cal BP), there is a distinct change to a light orange-brown silty-clay at the base of Unit 5. A shift to light-grey clay deposition began as early as 1.7 ka cal BP, the basal age of Unit 6, at 34 cm. After a return to olive-brown silt in Unit 7, (c. 1.2 ka cal BP), there is a shift to a light grey clayey-mud with silt and sand-rich layers (Unit 8: 20–12 cm depth; 1.2–0.5 ka cal BP). Unit 9 (12–5 cm; 0.78– -0.005 ka cal BP) is composed of laminated olive-brown silt, similar to Units 5 and 7. Unit 10 is a thick sub-aquatic moss composed entirely of *Drepanocladus l. sp.* AM of post-bomb age (1950-2011 CE) (Table 2). Lithofacies units and chronological descriptions for Kiteschee Lake are summarised in Figure 4 and Table 2.

*Diatom analysis:* The Kiteschsee Lake diatom record is characterised by short-term changes in diatom communities superimposed upon the steady ecological and environmental evolution of the lake during the mid-late Holocene. Although certain species found respond directly to climatic variables (temperature, precipitation and wind strength), the overall community structure response in Kiteschsee Lake was primarily associated with changes in lake-ice cover and turbidity, reflected in  $\beta$ -diversity, DCCA scores (Figure 4A).

Variations in key indicator species and the reconstructed chlorophyll-a were used to examine the impact of climate-driven glacier readvances in the mid-southern BIC foreland (Figure 4, 5). A notable shift in XRF-CS geochemistry exists at c. 60 cm depth (Figure S1), but we found no clear bio-geochemical evidence of (glacio) marine sedimentation. Although diatom concentration in the lowermost zone was too low for quantitative analysis, all sedimentary units younger than c. 6 ka cal BP (c. 63 cm) contained freshwater diatoms. The  $\delta^{13}\text{C}$  value from 73.5–74.5 cm depth is also consistent with a freshwater sedimentary environment, implying that it is younger than c. 6.5 ka cal BP. The increase in diatom accumulation rate from c. 1.9 ka cal BP is linked to the increased input of minerogenic sediments and dry mass accumulation

rate between c. 1.9 and 1.6 ka cal BP (Figure S8). Redundancy analysis (RDA) using key geochemical variables associated with tephra from XRF-CS analysis showed that tephra deposition did not have a significant impact on the diatom community composition in the Kiteschsee Lake record ( $p=0.596$ ) (Figure S9, S10).

Using constrained cluster analysis (Figure 4A, S9), we divided the diatom record from 47 cm upwards into five zones (Supplementary Materials for details). Diatom Zones 4 and 5 are of most interest to this study because they correspond to a decline in aerophilic and littoral diatom species and concurrent increases in species associated with more turbid conditions (e.g., *Stauriserella pinnata*), and lower DCCA scores (Figure 4A, 5B, S9). The transition to increased minerogenic input in lithofacies Unit 8 coincides broadly with the start of Diatom Zone 4 (22–13 cm: 1.3–0.8 ka cal BP). Trends initiated in Diatom Zone 4 continued in Zone 5 until ~5 cm depth (mid-late C20th onwards), when there was a significant shift to an aquatic moss (*Drepanocladus longifolius* (Mitt. Paris) sp.) dominated environment (Lithofacies Unit 10).

[insert Figure 5.]

## Discussion

We use new and published cosmogenic and radiocarbon ages, and lake sediment data, to assess the deglaciation and glacial readvance history of the BIC in the mid–late Holocene. The discussion is divided into four readvance stages, grouping the eight probability ‘gaps’ that represent possible readvances on the Fildes Peninsula and King George Island as follows: Stage 1: c. 7 ka readvance/standstill (7.4–6.6 ka cal BP); Stage 2: mid Holocene retreat followed by up to four possible minor readvances at 5.3–4.8 ka cal BP; 4.5–3.9 ka cal BP; 3.3–3.0 ka cal BP; 2.6–2.2 ka cal BP; Stage 3 readvance: 1.7–1.4 ka; Stage 4 readvances: 1.0–0.7 ka cal BP and <0.7 ka cal BP. We assess the main mechanisms for Holocene climate change on the AP identified in Bentley et al. (2009) (principally, solar insolation, westerly winds, the Southern Annular Mode (SAM)) and other potential drivers (e.g., volcanism; Cole-Dai et al., 2021; Mann et al., 2021) of centennial to millennial-centennial scale glacier variability on KGI.

### Early Holocene deglaciation c. 12–8 ka cal BP

Previously published lake sediment records show that the terrestrial deglaciation of the southern-mid part of the Fildes Peninsula from LGM limits dates to c. 12–9 ka cal BP (Mäusbacher et al., 1989; Watcham et al., 2011) (Figure 1), and was broadly coeval with other areas of the SSI (Palacios et al., 2020). New radiocarbon data from the BIC foreland and Potter Cove show the ice cap had thinned and retreated substantially during the early Holocene, with large parts of the Fildes Peninsula and Potter Peninsula becoming ice free by c. 8 ka cal BP. The BIC had also retreated from the central area of Fildes Peninsula by c. 8 ka (Figure 1, 6A, 7E). New chronological evidence from the BIC foreland and the Shetland I moraine in this study is consistent with the deglaciation of Mondsee Lake (Figure 2), 45 m a.s.l., ~1,500 m away from the present glacier margin at this time (Schmidt et al., 1990). Recession progressed towards the higher elevation areas in the eastern sector located, above the C20<sup>th</sup> equilibrium line altitude (ELA) of ~150 m a.s.l., into areas located in depressions next to the former ice limit, and triggered an increased rate of isostatic uplift and falling relative sea level (Figure 7E, F) (Watcham et al., 2011).

[insert Figure 6.]

[insert Figure 7.]

### ***Stage 1 readvance***

Evidence of the Stage 1 readvance and a marine transgression, when relative sea-level was ~15 m a.s.l. after 7.5–7.0 ka cal BP in Potter Cove (Strelin et al., 2014) is consistent with radiocarbon ages from Potter Peninsula constraining deglaciation and cosmogenic evidence from the Fildes Peninsula constraining retreat after a readvance of the BIC across NW foreland at c. 7 ka cal BP (Figure 3, 5A, 6A, 7E-G). Watcham et al. (2011) established an upper Holocene marine limit of ~16 m a.s.l. at c. 8.0 ka cal BP, with a short-lived fall and return to that level at 7.0 ka cal BP indicative of a standstill in glacier retreat or a minor readvance.

Data constraining RSL from the SSI are more complex than other locations across Antarctica, and several different RSL curves have been suggested for the SSI (Pallàs et al., 1997; Bentley et al., 2005; Hall, 2010; Roberts et al., 2011; Watcham et al., 2011; Simms et al., 2012; 2021; Johnson et al., 2021). Bentley et al. (2005) initially proposed that RSL declined from an undated early-mid Holocene marine limit of ~16–18 m a.s.l. in a discontinuous manner, interrupted at 14.5–16 m a.s.l. by a readvance between 5.8–3.0 ka cal BP). In contrast, and with

more data, Hall (2010) and Watcham et al. (2011) both suggested RSL fall was largely continuous after 6 ka cal BP. A readvance at c. 7 ka is consistent with evidence of marine-freshwater transitions in isolation basins between 14–16 m a.s.l. on the Fildes Peninsula (cf. Watcham et al., 2011).

Following the readvance at c. 7 ka cal BP, new cosmogenic exposure ages from erratics in this study show the NW marine platform close to the present icefront on the Fildes Peninsula was ice-free between 6.6–5.3 ka cal BP (Figure 7E, Table 1). We found no evidence for cosmogenic nuclide inheritance; therefore, these erratics had not been reworked or eroded from pre-Pleistocene formations on the SSI and were most likely glacially transported from the AP prior to (or during) the LGM, then fully reset by reworking immediately prior to deglaciation. Tiefersee Lake and Hochlandsee Lake, located ~800 and ~1500 m from the present-day ice cap, also became ice free (Figure 2, 3E) (Mäusbacher et al., 1989; Mäusbacher, 1991).

Marine sediments with bivalve shells dated to between 6.8 and 5.3 ka cal BP (95% probability phase range; Figure 5A), were reworked into till at the extreme east of the Shetland I moraines, during a subsequent readvance, most likely between 5.3–3.9 ka cal BP as RSL fell below 15 m a.s.l. (Figure 2, 6B, 7E; Table 2, S3). Moraines in this area were located close to the former sea-level prior to reworking, and Holocene raised beaches up to c. 16–18 m a.s.l. extended under the BIC in the Artigas Beach area may have survived relatively unmodified beneath the ice cap (John and Sugden, 1971; Hall, 2003).

Persistently warmer atmospheric temperatures across the Peninsula between 7 and 6 ka cal BP led to reduced sea-ice and warmer-than-Holocene-average sea-surface temperatures in marine sediment records from Maxwell Bay and the Palmer Deep between 8.2 and 5.9 ka cal BP (Milliken et al., 2009; Mulvaney et al., 2012; Etourneau et al., 2013). Overall, the marine-terminating BIC was less extensive during this shell colonization phase between 6.8–5.3 ka cal BP (Hall, 2007; Fretwell et al., 2010), and had retreated within its present-day limits by c. 6.0 ka cal BP, with a glacier front (re)occupying the inner bays of Collins Harbour in Maxwell Bay (Mäusbacher et al., 1989; Tatur et al., 1999; Yoon et al., 2006; Milliken et al., 2009; Simms et al., 2011; Watcham et al., 2011; Chu et al., 2017). Elsewhere on the Peninsula and around Antarctica ice sheets thinned and ice shelves retreated (Roberts et al., 2008; Hein et al., 2016; Verleyen et al., 2017; Johnson et al., 2020). Windier conditions at this time in the Southern

Ocean have been linked to rising global temperatures in the early to mid-Holocene (Marcott et al., 2013; Saunders et al., 2018; Kaufman et al., 2020) (Figure 5G, 7B, 7C).

### ***Stage 2 Readvances c. 5.3–2.2 ka cal BP***

There are four probability ‘gaps’ when minor readvances could have occurred between c. 5.3–4.8 ka cal BP, 4.5–3.9 ka cal BP, 3.3–3.0 ka cal BP and/or 2.6–2.2 ka cal BP, but evidence for glacier readvances on KGI and SSI is limited across the mid- to late Holocene transition. We found no direct evidence of BIC readvance on the Fildes Peninsula beyond its present-day limits between 5.3–3.9 ka cal BP, but it seems likely that the general warming transition into the mid Holocene Hypsithermal (MHH) (as defined by Bentley et al. (2009)) was interspersed by at least one minor/short-lived readvance (Figure 3D, 5E, 5H).

The Stage 2 readvance probability ‘gap’ between 5.3–4.8 ka cal BP (Figure 5A) overlaps with a hypothesised local readvance in the Stranger Point area of Potter Peninsula at c. 5 ka cal BP (Emslie et al., 2019), with persistently colder temperatures in the Yanou Lake record between 6.0–3.7 ka cal BP, and glacier readvances elsewhere on the NAP (Figure 5C, E, H). Surface temperatures decreased, and sea-ice cover increased in Maxwell Bay between c. 5.9 and 2.6 ka cal BP and summer sea-ice occupied the inner Collins Harbour of Maxwell Bay until at least c. 1.7 ka cal BP (Simms et al., 2011).

In contrast, on land, warmer-than-Holocene-average atmospheric temperatures across the SSI and AP and between c. 6.0 and 2.8 ka cal BP led to extensive moss banks and the formation of widespread lacustrine moss deposits on KGI (Figure 3E, 5F) (Björck, 1991a; Björck, 1991b; Sun et al., 2005; Hall, 2007; Bentley et al., 2009; Mulvaney et al., 2012; Foster et al., 2016; Yu et al., 2016; Roberts et al., 2017). Aquatic moss was continuously present between 6.0–4.5 ka cal BP (Figure 3E) (Milliken et al., 2009; Watcham et al., 2011; Roberts et al., 2017).

### ***Stage 3 readvance (late Holocene Neoglacial) c. 1.7–1.3 ka***

The BIC readvanced only by tens of metres relative to its present position in the Artigas Beach area in the second half of the late Holocene and most of the present-day Shetland I moraine was established during Stage 3 and/or Stage 4 readvances.

New mean radiocarbon ages from moss fragments incorporated into a small moraine on top of a raised beach at 7–6 m a.s.l. on Artigas Beach, indicate a two stage readvance of the BIC in

this area during the late Holocene: the first, after the formation of the raised beach and sometime between c. 1.7 and 1.4 ka cal BP (Stage 3 readvance) and a second at <0.7 ka cal BP (Stage 4 readvances). New radiocarbon ages also constrain the initial formation of the present-day Shetland I moraine from Valle Norte towards its eastern extension (see Figure 2 for location), where moraines are still ice-cored and in contact with the present ice. There, we found moraine crests several meters above the ice-level, similar to those found in Valle Norte (Hall, 2007), which post-date c. 1.2 ka cal BP, the age of the youngest new sample dated. This age means that the eastern part of the BIC was landward of its present limit, with mosses colonizing its foreland until at least 1.2 ka cal BP.

The Stage 3 readvance occurred after the start of the ‘Neoglacial’ on the AP, defined as  $2070 \pm 50$  cal BP by Čejka et al. (2020). The most significant changepoint in the Yanou Lake reconstructed temperature record at 1.3 ka cal BP coincides with a shift to colder temperatures and a return to more turbid conditions in Kiteschsee Lake (Figure 5B, C, S11). At this time, the Davies Heights, which is adjacent to Kiteschee Lake and close to the present-day ELA, could have sustained more permanent ice-cover, even with a relatively minor reduction in temperature (Figure 6C). This was broadly synchronous with two phases of glacier expansion on the north-eastern AP (Kaplan et al., 2020) (Figure 5E), in Southern South America between c. 1.5 and 1.1 ka cal BP (Strelin et al., 2014; Moreno et al., 2018), but during a predominantly warmer phase in composite ice core records from elsewhere in Antarctica (Masson et al., 2000).

#### ***Stage 4 readvance(s) c. 1.0–0.7 ka cal BP and after <0.7 ka cal BP***

Probability ‘gaps’ in the Shetland I moraine radiocarbon ages (Figure 3) suggest that readvances could have occurred at c. 0.7, 0.4 and/or <0.2 ka cal BP. Hall (2007) determined that glacier readvances of the Shetland I moraines occurred at Valle Norte and Valle Klotz (Figure 1 for locations) between c. 1.0 and 0.7 ka cal BP (Figure 2, 3C, 5A). The ice in these areas advanced in lobes, which spread out into ice-marginal depressions, forming several small thrust moraines, c. 300 m from the prominent moraine ridges (dated in this study). A final readvance bulldozed marine and fluvial sediments immediately in front of the present-day BIC limits into their current positions sometime after c. <0.2 ka cal BP (Figure 3B). Hall (2007) described this as the ‘last readvance’ and labelled it the most extensive phase of glacier readvance of the last c. 3,500 years. These ages are broadly aligned with persistently low DCCA Axis 1 ( $\beta$ -diversity) values in Diatom Zone 5 of Kiteschee Lake, reflecting resuspension

of sediment in a more-persistently turbid lake environment, which is indicative of reduced lake-ice, increased meltwater and/or (stronger) winds (Figure 4, 5B, S5F, S11).

The Stage 4 readvances, and increased lake turbidity, occurred during a phase of sustained colder reconstructed summer air temperatures in the nearby Yanou Lake record (Figure 5G, S11) (Roberts et al., 2017) after the Medieval Climate Anomaly (MCA) (c. 1.2 and 0.8 ka cal BP - Bentley et al. (2009); 1250–950 CE - Mann et al. (2009)). No evidence of substantial glacial readvance exists on KGI, northern AP region during the MCA (Charman et al., 2018; Emslie et al., 2019; Oliva et al., 2019; Kaplan et al., 2020), and marine sediment records from Maxwell Bay and Potter Cove imply no substantial increases in sea-ice at this time (Milliken et al., 2009; Monien et al., 2011). Increased fine-grained sedimentation in Potter Cove during the MCA has been linked to warmer conditions, including warmer sea surface temperatures from the western AP (Milliken et al., 2009; Hass et al., 2010; Watcham et al., 2011; Etourneau et al., 2013). Some Southern Hemisphere proxy records between 60 and 90°S (Figure 5K), and most records from north of 60 °S, show warmer conditions during the MCA (Marcott et al., 2013), but no sustained upturn in temperature exists in the James Ross Island ice core record or the Yanou Lake GDGT palaeotemperature record during the MCA (Figure 5G, S11) (Björck et al., 1993; Mulvaney et al., 2012; Roberts et al., 2017; Kaplan et al., 2020) (Fig 5I, S12).

The final <0.2 ka readvance was broadly synchronous with geomorphological and sedimentological evidence for glacier advance on Potter Peninsula/Cove (Hass et al., 2010; Heredia Barión et al., submitted), elsewhere on the SSI, the AP and several sub-Antarctic islands (Hall and Parry, 2004; Bentley et al., 2007; Hall, 2007; Bentley et al., 2009; Hall, 2010; Roberts et al., 2010; Davies et al., 2013; Davies et al., 2014; Guglielmin et al., 2016; van der Bilt et al., 2017; Simms et al., 2021; Verfaillie et al., 2021), extreme cold ‘events’ in most Southern Hemisphere palaeorecords (Neukom et al., 2014), and the Northern Hemisphere ‘Little Ice Age’ (0.55–0.25 ka cal BP, 1,400-1,700 CE) (Mann et al., 2009; Kaufman et al., 2020).

Evidence of Stage 3 and 4 readvance(s) and persistently colder conditions in lake records from the mid-outer areas of the Fildes Peninsula (Figure 5C) (Roberts et al., 2017) in the last 2,000 years mirrors a possible isostatic rebound event followed by an accelerated rate of RSL decline between 1.5–0.5 ka BP (Figure 7F) (Hall, 2010; Simms et al., 2012; Johnson et al., 2022) and raised beaches at <10 m a.s.l., which are thought to be contemporaneous with late Holocene



readvances on KGI and Livingston Island dated to between c. 0.5 and 0.25 ka cal BP (Sugden and John, 1973; Hall and Perry, 2004; Hall, 2010; Simms et al., 2012; Johnson et al., 2022). Restricted late Holocene readvances on the Fildes Peninsula are consistent with the limited late Holocene ice-loading scenario of the W12a GIA model (Whitehouse et al., 2012a; Whitehouse et al., 2012b). Greater ‘Neoglacial’ ice-loading is broadly supported by field data, implying more substantial readvances and/or variable ice-mass (un)loading elsewhere on the SSI during the late Holocene.

### ***Mechanisms of Holocene glacier fluctuations***

Deglaciation of the Fildes Peninsula and other peninsulas on KGI to present day limits occurred in two broad phases during the early-mid Holocene. The first phase between c. 15 and 8 ka cal BP occurred during a period of rising spring/summer insolation at 62 °S and relatively elevated global solar irradiance in the transition into the present (Holocene) interglacial (Figure 7A, D, E) (Steinhilber et al., 2009; Baggenstos et al., 2019; Reynhout et al., 2019). This coincided with a period of sustained warmth and persistently positive Southern Annular Mode (SAM)-like conditions (Figure 7D). Meanwhile, data from the BIC foreland moraines and other readvance constraints across KGI suggest that ice caps on the SSI likely responded to millennial- to centennial-scale changes in solar forcing (insolation and irradiance), westerly wind strength and the SAM in the mid-late Holocene.

The SAM is the primary mode of annual to centennial scale variability in atmospheric circulation in the Southern Ocean (Marshall, 2007) and reflects the zonal mean sea level pressure difference between Antarctica (>65°S) and the mid latitudes (40°S); hence, the SAM index is, in effect, a measure of the longitudinal mean SHW strength (Marshall, 2007). Positive SAM phases are associated with lower pressure at higher latitudes, more poleward focussed SHW, with stronger Westerlies in summer and stronger Easterlies in winter over the NAP and SSI (Martin et al., 2021). During the Holocene, stronger and more poleward SHW are thought to be the centennial- to millennial-scale expression of decadal-centennial warming and positive SAM-like conditions (Charman et al., 2018; Moreno et al., 2018; Perren et al., 2020). Conversely, phases of enhanced El Niño drive warmer interannual global mean temperature variability and negative phases of the SAM, which are characterised by colder and more humid conditions on the NAP, SSI and sub-Antarctic islands, and have increased over the last 4 kyr (Figure 7C) (Wang and Cai, 2013; Turner et al., 2016; Reynhout et al., 2019; Kaplan et al., 2020; Verfaillie et al., 2021).

Since 1957 CE, the longer-term pattern of interannual variability in the SAM has been altered by the ozone hole over Antarctica, which, combined with increased greenhouse gases and global temperature, has led to more positive SAM-like conditions, similar to the early Holocene, and enhanced the ENSO (Marshall, 2007) (Figure 7D). Increased cyclonic activity in the Drake Passage between 1996–2005 CE has been linked to a more poleward focussed SHW and more positive SAM, driving sea-ice poleward and increasing the advection of warm air across the NAP and SSI during winter (Olivia et al., 2017; Marshall et al., 2017). Between 2006–2015 CE colder winters pushed sea ice north, increasing snowfall across the NAP and doubling snow cover on some SSI islands between 2009–14 CE (Olivia et al., 2017; de Pablo et al., 2017). In addition, the SAM and a deepening Amundsen Sea low have been interacting in recent decades, bringing warmer air from the Atlantic and north-easterly winds with cooler air from the Weddell Sea (Liu et al., 2005; Davies et al., 2014; Kaplan et al., 2020). As a result, the NAP and SSI became colder and wetter, and its glaciers have gained mass (Thomas et al., 2008; Goodwin et al., 2016; Oliva et al., 2017; Falk et al., 2018).

Accordingly, longer-term changes in temperature and precipitation that drove glacier readvances on the NAP and SSI during the Holocene were most likely the result of annual to millennial scale variations in the position and intensity of the SHW, modulated by changes in the SAM (Moy et al., 2008; Bentley et al., 2009; Lamy et al., 2010; Varma et al., 2012).

The Stage 1 readvance occurred during a downward trend in global solar irradiance between c. 7.5 and 7.0 ka cal BP which coincided with the transition to colder and more negative SAM-like conditions as the SHW migrated marginally equatorward (Figure 7A–C, G). Glacier readvance(s) on KGI Glaciers on KGI are regionally coherent with readvances on the eastern AP (Kaplan et al., 2020; Figure 5E) and in Patagonia at c. 7 ka (Strelin et al., 2014; Reynhout et al., 2019), implying a common forcing mechanism.

Following the Stage 1 readvance, increasing spring/summer insolation over the SSI and global solar irradiance and weaker/ marginally equatorward westerlies reduced storminess drove glacier retreat on the Fildes Peninsula and the Potter Peninsula between c. 6.6–5.3 ka cal BP. (Figure 7A–D, G, H) (Bentley et al., 2009; Quade and Kaplan, 2017; Saunders et al., 2018; Reynhout et al., 2019; Heredia Barión et al., submitted). Retreat on KGI coincided sustained temperature around the 0–12 ka mean in the Southern Hemisphere high latitudes (Marcott et

al., 2013) and with the ‘Holocene global thermal maximum’, centred on 6.5–6.0 ka cal BP (Figure 5G). Although the latter is more clearly defined in Northern Hemisphere proxy records, average global temperatures were 0.7 °C warmer than the pre-industrial 19th Century era (Kaufman et al., 2020).

While transitions into and/or out of short-lived phases of marginally more positive SAM-like conditions in the mid-late Holocene could have increased snow accumulation on the SSI, the combination of generally colder and more negative SAM-like conditions, stronger westerly winds, and an enhanced ENSO from c. 4 ka onwards (Garreaud, 2007; Yang et al., 2019) led to deteriorating regional climatic conditions that lowered the ELA across the SSI and drove late Holocene glacier readvances on KGI (Figure 5F-G, 7A-D, G, H, S12).

Late Holocene (<4.2 ka cal BP) Stage 2, 3, and 4 readvances occurred as spring/summer insolation at 62°S declined and around global irradiance minima. These readvances are anti-phased with aquatic moss layers in lake records from outer Fildes Peninsula lakes unaffected by glaciofluvial inputs. Generally stronger westerly winds and colder/more negative SAM-like conditions likely drove marginal increases in snow accumulation and greater turbidity in Kiteschsee Lake (Figure 4, 5B, 7B), but an equatorward shift in the core westerly wind belt storm belt to <55°S likely reduced storminess (and precipitation), restricting readvances to around or within present day limits (Figure 4, 5B, 7B, 7C).

From the mid-late C20th onwards, the mode of sedimentation in Kiteschsee Lake, other lakes on the mid-outer Fildes Peninsula and lakes elsewhere on King George Island that are now detached from direct glaciofluvial influence has shifted dramatically to one dominated by sub-aquatic moss (Figure 4A, S1). We link this dramatic ecological shift to the Recent Rapid Regional warming (RRR) on the northern AP and SSI (Bentley et al., 2009) and increasingly positive SAM-like conditions drawing stronger SHW poleward in the late C20th and early C21st (Figure 7D) associated with anthropogenic activity (e.g., the ozone hole, increasing greenhouse gas emissions) that are drawing

## **Conclusions**

We reconstructed the mid–late Holocene deglaciation and readvance history of the Fildes Peninsula, King George Island/Isla 25 de Mayo (KGI), South Shetland Islands, NW Antarctic Peninsula, using new cosmogenic exposure ages from glacial erratics on the foreland of the

Bellingshausen Ice Cap (BIC), radiocarbon dating of terrestrial mosses and marine *Laternula* sp. bivalve shells found in moraines on the Artigas Beach and Valle Norte sectors of BIC foreland.

Using a novel statistical approach, we developed a new deglaciation and four-stage Holocene readvance model for the BIC on the Fildes Peninsula

Deglaciation on KGI began after c. 15 ka and had progressed to within present-day limits on the Fildes Peninsula, its largest ice-free peninsula between c. 6.6–5.3 ka cal BP as spring/summer insolation at 62°S gradually increased, leading to generally warmer and more positive SAM-like conditions during the early Holocene.

Probability density phase analysis of new (n=25) and existing chronological data constraining Holocene glacier advance on the Fildes Peninsula and other ice-free peninsulas of KGI (n=80 in total) revealed up to eight 95% probability ‘gaps’ when readvances could have occurred, which we grouped into four stages:

- Stage 1: a readvance and marine transgression, well-constrained by field data, between c. 7.4–6.6 ka cal BP.
- Stage 2: four less well-constrained readvance probability ‘gaps’ between c. 5.3–4.8 ka cal BP, 4.5–3.9 ka cal BP, 3.3–3.0 ka cal BP and 2.6–2.2 ka cal BP
- Stage 3: a well-constrained readvance between c. 1.7–1.5 ka when terrestrial mosses were embedded in the Shetland I moraine and the BIC extended landward of its present position on its eastern flank
- Stage 4: two further readvances, one less well-constrained by field data between c. 1.3–0.7 ka cal BP (68% probability) when moss fragments and mid Holocene marine sediments were translocated into the moraines surrounding the present-day BIC and a prolonged phase of elevated turbidity existed in Kiteschsee Lake, and a well-constrained ‘final’ readvance after <0.7 ka cal BP, consistent with the ‘last readvance’ on James Ross Island, NE Antarctic Peninsula.

Similarities in the timing of mid-late Holocene glacier readvances on the SSI, the northern Antarctic Peninsula, some sub-Antarctic Islands and in southern South America imply underlying regional-hemispheric drivers were responsible.

The Stage 1 readvance occurred at the Holocene global solar irradiance minimum as colder/more humid (negative) Southern Annular Mode (SAM)-like and stormier conditions developed, with marginally stronger (poleward shifted) westerly winds over the SSI.

Readvances were more frequent after c. 5.3 ka cal BP, driven by reducing spring/summer insolation at 62°S and more persistent negative SAM-like conditions. However, weaker (equatorward shifted) westerlies and reduced storminess over the SSI restricted readvances close to present day limits.

Late Holocene readvances occurred around global solar irradiance minima and are anti-phased with aquatic moss layers in lake records from outer Fildes Peninsula lakes unaffected by glaciofluvial inputs.

Recovery from late Holocene ‘Neoglacial’ conditions in Kiteschsee Lake and other mid-outer lakes on the Fildes Peninsula in the post-bomb era (>1950 CE) is characterised by the recolonisation of well-developed aquatic moss at the lake sediment-water interface. This was driven by late 20th Century/early 21st Century warming of the northern Antarctic Peninsula and more positive SAM-like conditions on KGI.

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## **Author Contributions**

PHB and SR designed the research and wrote the manuscript, constructed figures and tables with assistance from JD, IG, MB, GK; PHB, SR, EP, MB, DH, EH undertook fieldwork/sample collection. PHB and SB produced cosmogenic nuclide dating data; while CS provided <sup>10</sup>Be exposure dating analytical facilities; PHB, LW, GK, SR, MB undertook radiocarbon dating. Kiteschsee Lake core data and interpretations were undertaken by SR, EP, SD (ITRAX-XRF), JD, VJ (diatom analysis) IG, SB, RM, CH (tephra counting, geochemical identification of tephra layers); PHB, SR, CS, SB, GK, EP, MB, JS edited initial drafts and all authors read and commented on the final version. The authors declare no competing interests.

## **Data availability**

Key datasets generated in this study are included in the main text and Supplementary

Materials: <https://doi.org/10.6084/m9.figshare.21352569.v2>

All code, data, packages and package references, can be found at:

[https://github.com/stever60/Fildes\\_Peninsula](https://github.com/stever60/Fildes_Peninsula). Data can be obtained from the NERC EDS UK Polar Data Centre (PDC) as follows (PDC doi assignments pending): Chronological data:

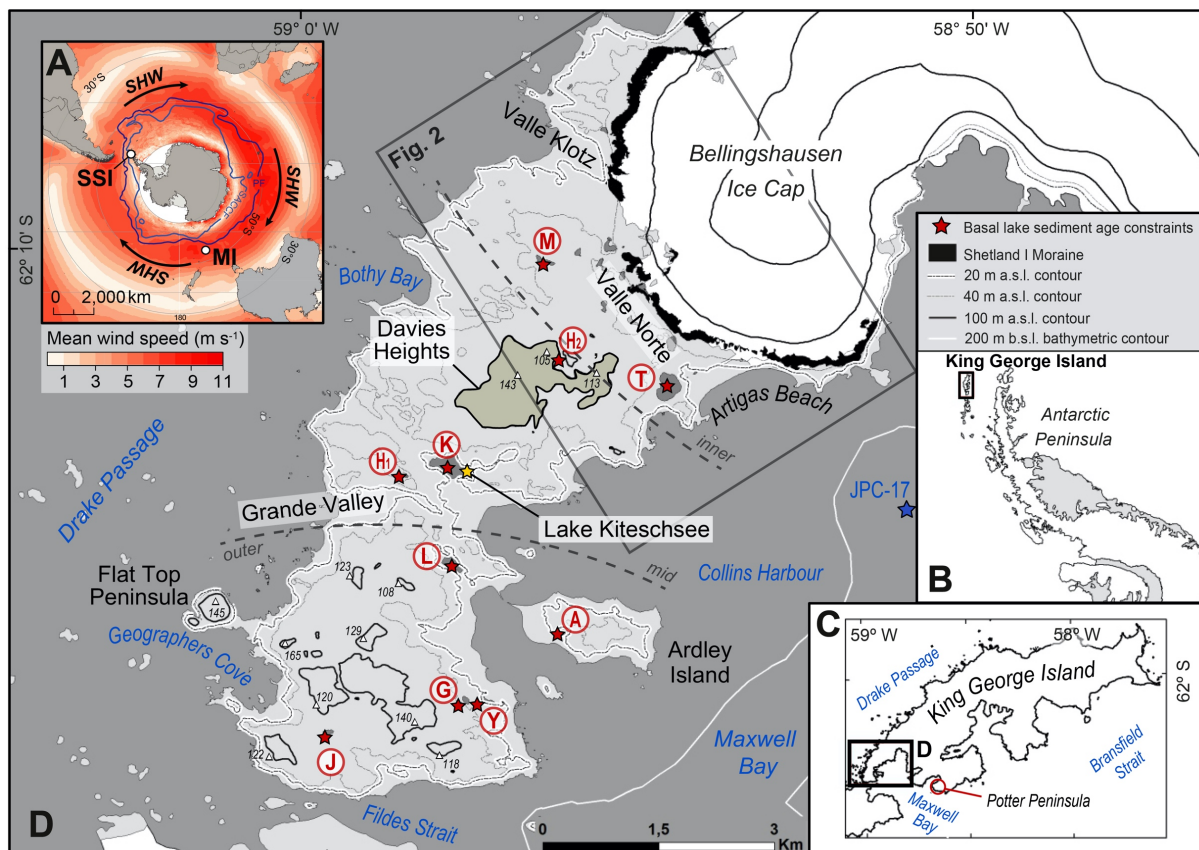
[https://github.com/stever60/Fildes\\_Peninsula](https://github.com/stever60/Fildes_Peninsula);

Lake sediment data: [https://github.com/stever60/Fildes\\_Peninsula](https://github.com/stever60/Fildes_Peninsula);

Advance/retreat data: [https://github.com/stever60/Fildes\\_Peninsula](https://github.com/stever60/Fildes_Peninsula)

Data are available on request from the Polar Data Centre ([polardatacentre@bas.ac.uk](mailto:polardatacentre@bas.ac.uk)), PHB ([herediabaronpablo@yahoo.com.ar](mailto:herediabaronpablo@yahoo.com.ar)) and SR ([sjro@bas.ac.uk](mailto:sjro@bas.ac.uk)). Final figure layouts and additional annotations were produced using Adobe Illustrator CC v. 2021 and CorelDRAW v. 2020.

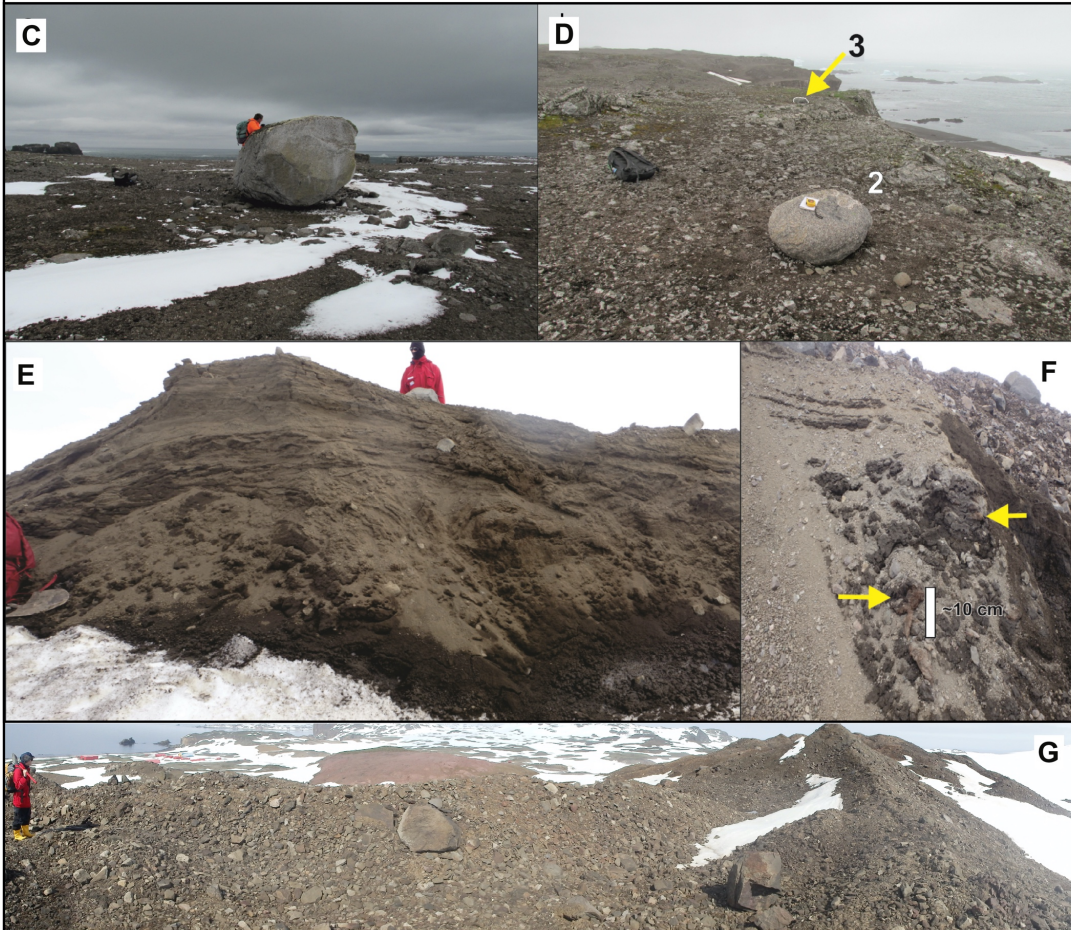
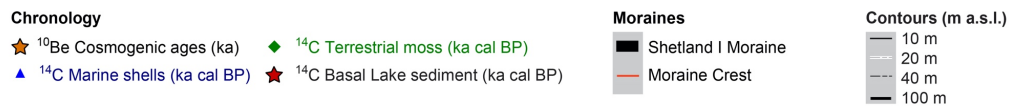
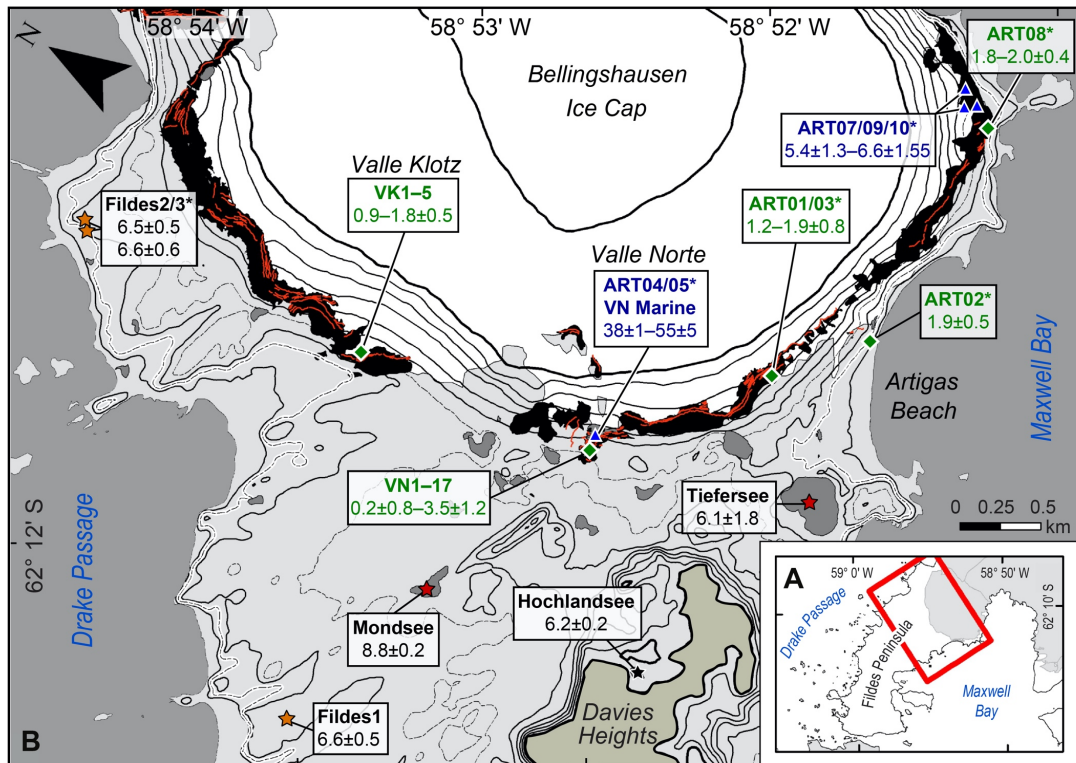
## Figures



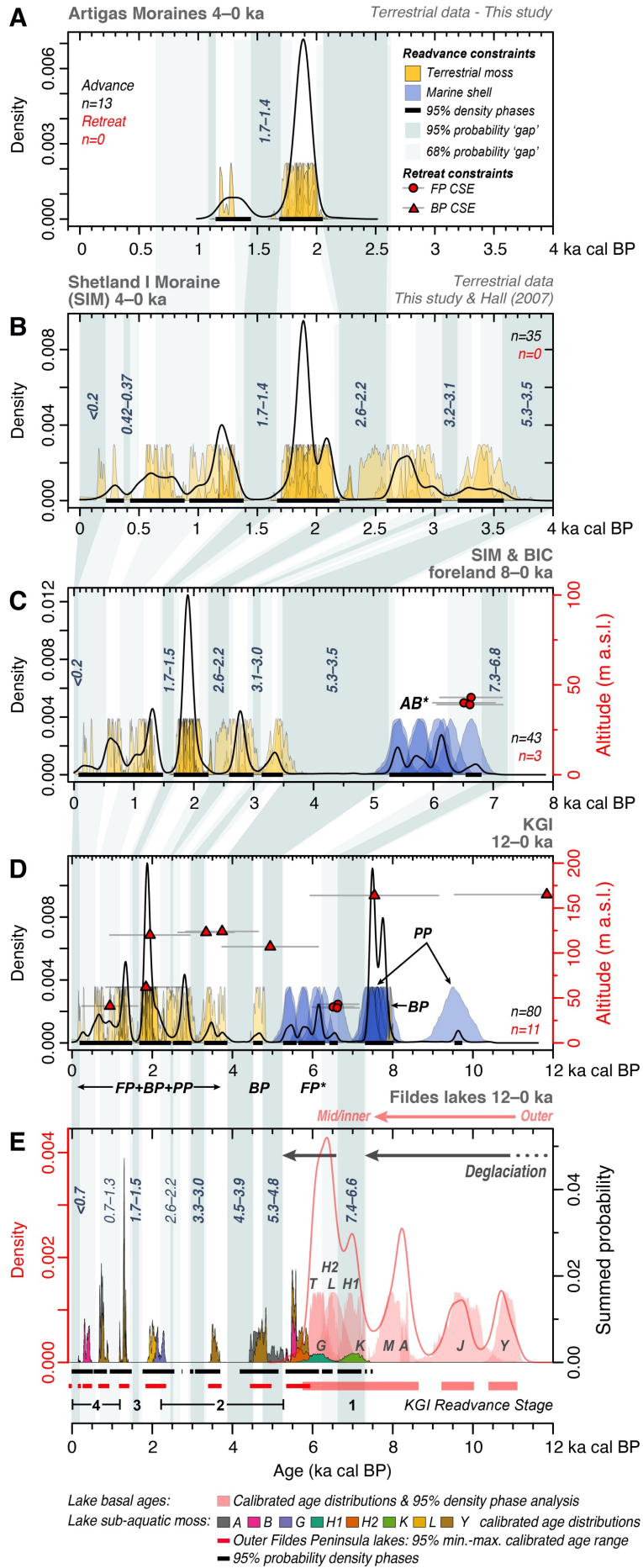
**Figure 1.** Geographic location of the study area. A) Location of the South Shetland Islands (SSI), currently south of the core Southern Hemisphere westerly Wind (SHW) belt. MI is Macquarie Island, located within the core SHW belt. B) Location of King George Island (KGI), South Shetland Islands on the north-western Antarctic Peninsula. C) KGI is framed by the Drake Passage and the Bransfield Strait. Box marks the position of the Fildes Peninsula. D) Map of the Fildes Peninsula showing minimum ages (a cal BP) for the onset of lake and offshore sedimentation. The Fildes Strait separates the Fildes Peninsula from Nelson Island, with Maxwell Bay to the east, and the open waters of the Drake Passage on its west coast. Lake record basal deglaciation ages (a cal BP  $\pm 2\sigma$ ), updated to Sh20/Marine20 for this study are as follows: Outer: J-Jurasee Lake (9,750 $\pm$ 780 a cal BP) (Mäusbacher et al., 1989); Y-Yanou Lake

(10,830±340 a cal BP) (Watcham et al., 2011; Roberts et al. 2017; G-Gaoshan Lake (7,250±120 a cal BP) (Watcham et al., 2011); A-Ardley Lake (8,310±120 a cal BP) (Watcham et al., 2011; Roberts et al., 2017); L-Long Lake (>6,480±180 a cal BP) (Yoon et al., 2006; Watcham et al., 2011); Mid: Marine core JPC-17 (5,800 a cal BP) (Simms et al., 2011); K-Kiteschsee Lake (>7,300±120 a cal BP) (this study and Mäusbacher, 1991); H<sub>1</sub>-Hotel Lake (6,960±150 a cal BP) (Tatur et al., 1999); H<sub>2</sub>-Hochlandsee Lake (6,150±400 a cal BP) (Mäusbacher, 1991); Inner: T-Tiefersee (Profound) Lake (6,120±360 a cal BP); M-Mondsee Lake (8,000±480 a cal BP) (Schmidt et al., 1990). Contour lines: Antarctic Digital Database (<http://www.add.scar.org/home/add7>); bathymetry: GEBCO (Weatherall et al., 2015); spot heights (white triangles): Servicio Geográfico Militar de Uruguay.





**Figure 2.** A–B) Map showing the position of cosmogenic  $^{10}\text{Be}$  exposure ages of sampled erratics (orange stars) and calibrated radiocarbon ages of organic remains in the Shetland I moraine (black shading) and minimum ages of the onset of lake sedimentation (red stars). Inset map shows the area where the study sites are located (red box) BIC is Bellingshausen Ice Cap. VK = Valle Klotz, VN = Valle Norte, both sampled and dated by Hall (2007); asterisks indicate new data in this study. C–F) Field photographs showing the erratics sampled on the NW marine platform for terrestrial cosmogenic nuclide exposure age dating and terrestrial sediments preserved within the Shetland I moraine. C) Fildes 1 and D) Fildes 2 and 3 boulders sampled for cosmogenic analysis. Boulder dimensions and locations are in Table 1. E & F) Radiocarbon dating samples from locations ART-01 and ART-03 embedded in moraine sediments (Table 2). G) View from the crest of the Shetland I moraine in the Artigas Beach area looking approximately south towards the Uruguayan Station.

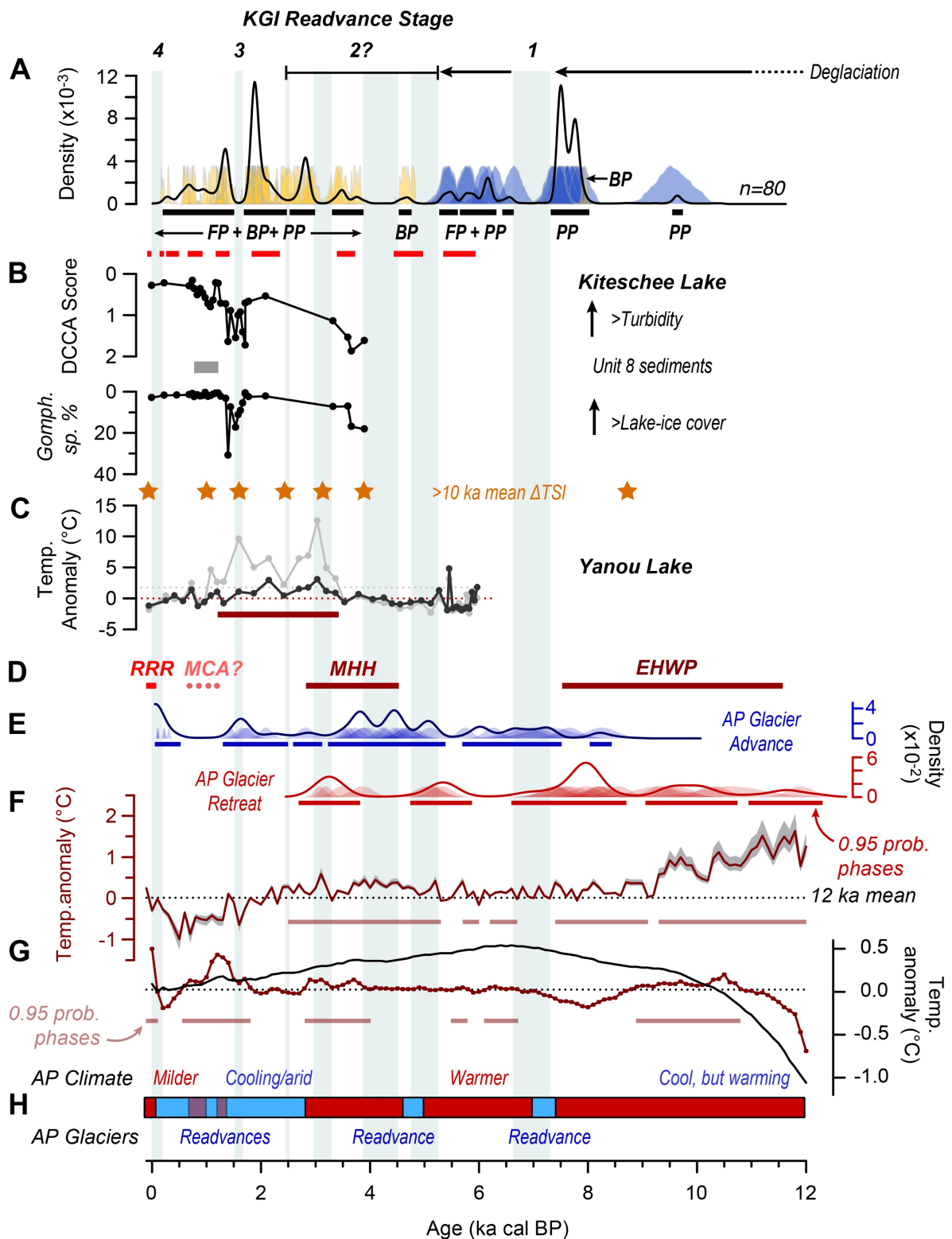


**Figure 3.** A) Radiocarbon ages and probability density phase model (black bars 95%) for terrestrial mosses dated in this study from the Artigas Beach sector moraines between 4–0 ka cal BP (i.e., sample locations ART01–03 only in Figure 2; maximum age constraints on readvance). Dark and light grey shaded zones, with 95% probability timings, indicate periods when glacier readvances could have occurred (95% and 68% probability, respectively). B) Combined radiocarbon ages and probability density phase model for terrestrial mosses dated in this study from the Artigas Beach sector moraines and the Valle Norte–Valle Kotz sector moraines dated by Hall (2007) shown in Figure 2A; all data are maximum age constraints on readvance. C) Radiocarbon ages and probability density phase model for the Shetland I moraine 8–0 ka cal BP, consisting of terrestrial mosses from this study and Hall (2007) and marine shells dated from the Artigas Beach (AB\*) sector moraines in this study. Cosmogenic surface exposure (CSE) ages from the BIC foreland from this study are also shown (red circles; minimum age constraints on retreat). D) Probability density phase model for all maximum age constraints on glacier readvance on King George Island between 12–0 ka cal BP. Cosmogenic nuclide exposure ages from this study (red circles) and Barton Peninsula (Seong et al., 2008) (red triangles) provide minimum constraints on retreat. BP=Barton Peninsula; FP=Fildes Peninsula, PP=Potter Peninsula. E) Probability density phase model for basal deglaciation ages (red line and bars) compared to a summed probability density profile of radiocarbon ages from aquatic moss layers in lake sediment records from the Fildes Peninsula. King George Island (KGI) Readvance Stages 1–4 are also shown. Aquatic moss deposition represents lacustrine conditions similar to the present day. Lakes are as shown in Figure 1. The units for the probability density phase analysis are  $\text{ka}^{-1}$ ; data sources are described in code input datafiles.





**Figure 4.** Summary of key stratigraphic, diatom and chronological data from the Kiteschsee Lake record; I = Tephra layers in the Kiteschsee Lake record, II = ITRAX XRF-CS optical image, III = ITRAX X-radiograph (positive image), IV = Summary lithology and RGB Munsell colour. B) Bayesian best fit age-depth model for the Kiteschsee Lake record. Sample thicknesses were 0.5 cm or 1 cm, and no zero-depth surface age was set in the age-depth model. Run settings and priors are shown in Figure S4. Modelled ages and errors were derived from the MCMC ‘best-fit’ weighted mean ages from the BACON age-depth model. Initial age model runs were fine-tuned by tephrastratigraphic correlation to the nearby Yanou Lake record and the inclusion of radiocarbon ages from aquatic moss immediately below and above the ‘T5’, between the ‘T3a’ and ‘T3b’ tephra layers, and immediately below the ‘T1a’ layer in the Yanou Lake record (Supplementary Material for details). Radiocarbon ages from the 77 cm long, c. 7,600-year-old sediment record from Kiteschsee Lake are in chronological order, except for a calibrated age of  $9,900 \pm 240$  a cal BP near the base of the fine grey (glaciofluvial) silt and clay deposits of Unit 8 at 21–22 cm depth (Table 2). Bulk sediments at 7–8 cm, 54.5–55 cm, 64.5–65 cm depth could not be dated due to insufficient carbon. In summary, the Kiteschsee Lake record has a basal age of 7,470 a cal BP at 77 cm depth (weighted mean modelled age) (7,820–7,110 cal a BP min.–max. 95% confidence age range; 7,480 cal a BP median age). The mean age-depth model 95% confidence age range is 470 years, with a minimum 95% confidence age range of 4 years at 4.5 cm, in the post-bomb era and a maximum 95% confidence age range of 1,230 years at 51.5 cm; 83% of the measured ages overlap with the age-depth model (95% ranges). Elevated dry mass accumulation rates and increased sand deposition within Unit 6 and the glaciogenic Unit 8 are indicative of increased erosion and meltwater input from c. 1.7 ka cal BP onwards (Figure S3, S8). The radiocarbon age of  $9,380 \pm 30$   $^{14}\text{C}$  years from 21–22 cm depth at the base of Unit 8 had a heavily depleted  $\delta^{13}\text{C}$  value of  $-32.6$  ppm, likely due to reworking of old (possibly marine) carbon deposited in/around the Davies Heights massif. The uppermost sediments, immediately below the living aquatic moss, have a measured radiocarbon age of  $1,350 \pm 30$   $^{14}\text{C}$  years ( $1.2 \pm 0.8$  ka cal BP) implying bulk sediment radiocarbon ages obtained from Unit 9 are either significantly less influenced by old carbon, that sedimentation rates declined, or a hiatus exists between Units 9 and 10. K1–8 are radiocarbon ages shown in Table 2, T1a, T3a, T5 and T7 are correlation ages of tephtras in the Yanou Lake record (Table S2). C) Age depth model and tephrastratigraphic correlation for the Yanou Lake record. D) An updated Bayesian age-depth model (2020 calibration) for the Yanou Lake record (cf. Roberts et al. 2017). All measured radiocarbon ages were included in age-depth models and the SHCal20 calibration curve used (Hogg et al., 2020).



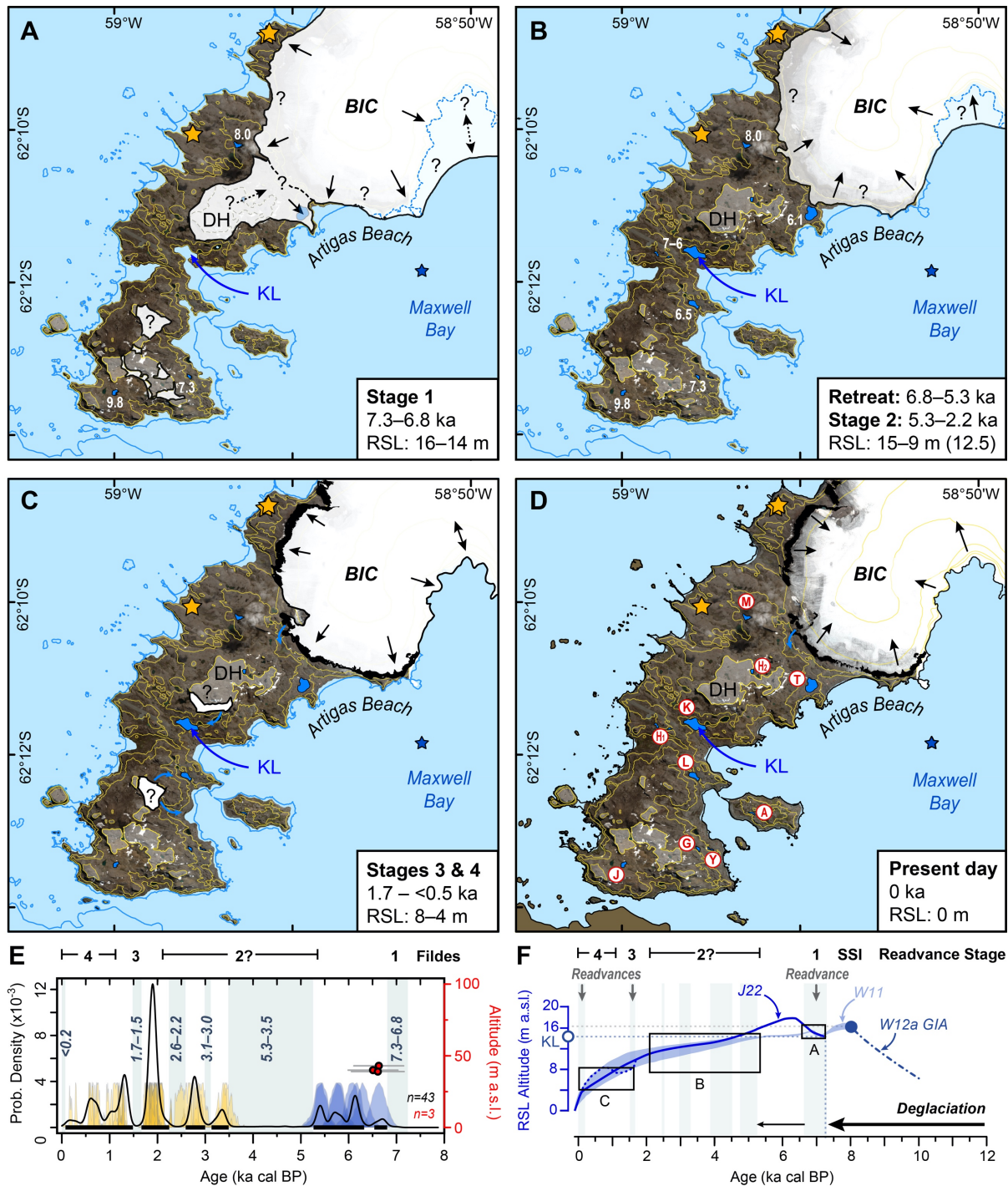
**KGI readvances: max. age constraints**

- Terrestrial moss  $^{14}\text{C}$  ages
- Marine shell  $^{14}\text{C}$  ages
- 95% probability 'gap'
- Outer Fildes Peninsula lakes: Sub-aquatic moss 95% min.-max. age range
- Density phases (95% prob.)

**Figure 5.** Regional palaeoclimatic context for Holocene glacier fluctuations on the South Shetland Islands, Antarctic Peninsula. A) Probability density phase model for glacier

readvances Stages 1–4 on King George Island between 12–0 ka cal BP (Figure 3D for details); BP=Barton Peninsula; FP=Fildes Peninsula, PP=Potter Peninsula. B) Diatom-inferred reconstructions and key indicator diatom species for Kiteschsee Lake. Lower DCCA scores reflect more turbid conditions, and lower *Gomphonema* sp. percentages represent less aerophilic conditions, and longer seasonal ice-cover. C) Glycerol dialkyl glycerol tetraether (GDGT)–mean summer air temperature (MSAT) anomaly reconstruction (light grey) from Yanou Lake (Foster et al., 2016; Roberts et al., 2017) (RMSE = 2.45 °C, not shown for clarity); the dashed line is the 6 ka mean MSAT anomaly for the dataset published in Roberts et al. (2017); the black line and datapoints represent a revised (interim) MSAT anomaly reconstruction for Yanou Lake (see Supplementary Materials for details). This and 6 kyr mean $\pm 1\sigma$  MSAT<sub>anomaly</sub> of 0.06 $\pm 1.50$  °C are more realistic than the previously published MSAT-GDGT data; orange stars are  $\Delta$ TSI peaks (see Figure 7). D) Climate syntheses for the Antarctic Peninsula (Ingólfsson et al., 2003; Bentley et al., 2009). E) Probability density phase model for glacier advance (blue) and retreat (red) data on the northern Antarctic Peninsula in Kaplan et al. (2020). F) Temperature anomaly data from the James Ross Island (JRI) ice core record, NE Antarctic Peninsula (Mulvaney et al., 2012); grey shading is the as published  $\pm$ error. G) Holocene palaeotemperature compilation showing the median temperature anomaly for 60–90 °S stack (dark red line) and the global temperature stack anomaly (black line) (Kaufman et al., 2020). The grey dashed line is the average median 12-kyr value for 60–90 °S temperature proxy data (0.02); errors in both datasets are mostly out of range and not plotted for clarity; the units for the probability density phase analysis are ka<sup>-1</sup>.

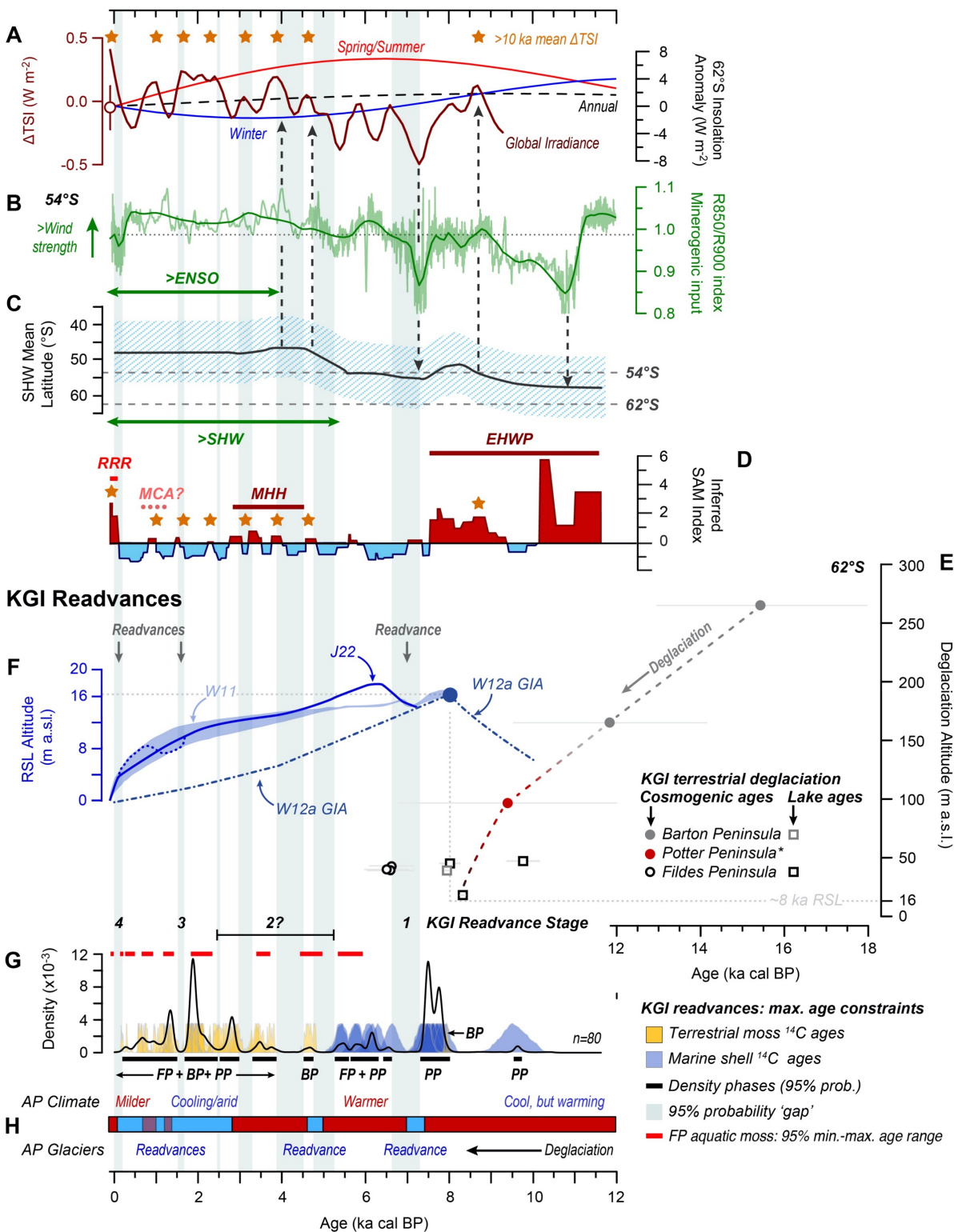




**Figure 6.** A–D) Schematic reconstruction of mid–late Holocene glacier readvance extents during Stages 1–4 on the Fildes Peninsula. Glacier extents shown in A–C are based on a compilation of new and published chronological data (panel E) for the Shetland I moraine (dark shading in C and D). E) Radiocarbon ages and probability density phase model for age constraints on glacier readvance on the Fildes Peninsula (n=43; Figure 3C for details; yellow distributions are radiocarbon ages of terrestrial material embedded in moraines; blue distributions are radiocarbon ages of marine material embedded in moraines; red circles with

grey external error bars are  $^{10}\text{Be}$  cosmogenic nuclide exposure ages obtained from boulders (orange stars in A–D); the units for the probability density phase analysis are  $\text{ka}^{-1}$ . F) Summarised relative sea level (RSL) curves for the South Shetland Islands (Watcham et al., 2011 – W11; Whitehouse et al., 2012a, b – W12aGIA; Johnson et al., 2022 – J22) used to determine the elevation of the coastlines in panels A–C; BIC = Bellingshausen Ice Cap; DH = Davies Heights; KL = Kiteschee Lake, which is a marine embayment in A; the blue star is the location of Maxwell Bay marine records referred to in the text (Milliken et al., 2009; Simms et al., 2011); white areas (90% transparency) bordered by solid black lines in A–C are schematic representations of BIC extent on the Fildes Peninsula for each stage; present-day moraines mapped in this study are shown in D; black arrows represent schematic glacier flow directions, except in B where they represent retreat close to (or within) present day limits between 6.8–5.3 ka BP, followed by minor readvances close to present limits between 5.3–2.2 ka BP; blue arrows are main meltwater pathways. Question marks in A and B indicate tentative locations of potential glacier fronts above 100 m a.s.l.; lakes sediment basal and/or freshwater transition ages shown as white text are in ka BP, labelled as in Figure 1. Black shaded areas around the BIC in C and D are the last readvance moraines, as shown in Figure 2. The 2013 CE satellite image of the BIC is shown in D, with studied lakes in blue (labelled as in Figure 1: white circles/red borders). Yellow lines are contours at 20 m, 40 m and 100 m elevation, with areas >100 m shaded white (30% transparency). The VHR satellite image in A–D is from Maxar Products: WorldView-2 image acquired 11th April 2013 © 2021 Maxar Technologies. This figure includes material copyright of Maxar Technologies, All Rights Reserved, reproduced here at low resolution with permission under a NERC-BAS educational licence and not included as part of a Creative Commons licence.

### Mechanisms of change



**Figure 7** Mechanisms of deglaciation and glacier fluctuations on the Fildes Peninsula and King George Island (KGI). A) Total Solar Irradiance, which is the power per unit area received from the Sun as measured, or reconstructed, on Earth –  $\Delta$ TSI represents the deviation from its present-day value (Steinhilber et al., 2009) and is compared with solar insolation received at

62°S during the Holocene (11.75 ka) (Laskar et al., 2004); annual insolation = dotted black line, austral spring/summer (SONDJF) insolation = red line; winter (JJA) insolation = blue line; the open dark red circle is the mean $\pm$ 1 $\sigma$   $\Delta$ TSI value for the last 10 ka; orange stars mark  $\Delta$ TSI peaks greater than the 10 ka mean value. B) Ultra-high resolution (69  $\mu$ m) hyperspectral (SPECIM) R850/R900 data, a proxy for mineral input into the Emerald Lake, Macquarie Island at 54°S (Saunders et al., 2018), which reflects changes Southern Hemisphere Westerly wind (SHW) strength during the mid-late Holocene; the dark grey horizontal dotted line is the R850/R900 dataset mean; dark green line is 100-year interval LOESS smoothing of the R850/R900. C) Hypothetical representation of changes in the mean annual latitudinal position of the core SHW intensity belt (dark grey line) and approximate 1 $\sigma$  latitudinal range of enhanced precipitation (light blue stipple) (Ariztegui et al., 2010; Quade and Kaplan, 2017). In reality, the SHW are more intense and focused during positive (warmer) phases of the SAM, but weaker, latitudinally broader and less focused during negative (colder) phases of the SAM, show in D. The light grey horizontal dashes are 54°S (Macquarie Island) and 62°S (South Shetland Islands); vertical dashed lines and arrows highlight the key hypothesised latitudinal shifts of the Holocene and their relationship to the global irradiance profile shown in A; ENSO = El Niño Southern Oscillation. D) Inferred Holocene SAM-index variability between positive (red) to negative (blue) SAM-like states based on Northern Arboreal Pollen data from Southern Chile (Moreno et al., 2018; Reynhout et al., 2019) (adapted from Kaplan et al. (2020)); EHWP = Early Holocene Warm Period; MHH = Mid Holocene Hypsithermal; MCA = Medieval Climate Anomaly; RRR = Recent Rapid Regional warming, from Bentley et al. (2009). E) Summarised relative sea level (RSL) envelope for the SSI (W11; light blue shading; Watcham et al., 2011) compared with the RSL curve in Johnson et al. (2022) (J22; dark solid and dotted blue lines) and the W12a GIA model (Whitehouse et al., 2012a, b); see original references for further details and data. F) Cosmogenic nuclide exposure ages and key lake basal age constraints on deglaciation for terrestrial ice-free peninsulas on KGI at different altitudes from the Fildes Peninsula (this study), the Potter Peninsula (Heredia Barión et al., submitted), and the Barton Peninsula (Seong et al., 2008; Oliva et al., 2019) (secondary plot); \* = new data (this study). G) Probability density phase model for glacier readvances Stages 1–4 on King George Island between 12–0 ka cal BP (Figure 3D for details); BP=Barton Peninsula; FP=Fildes Peninsula, PP=Potter Peninsula. H) Glacier-climate syntheses for the Antarctic Peninsula (adapted from Ingólfsson et al. (2003) and Bentley et al. (2009)).

Sample ID	AMS ID	Latitude (°S) Longitude (°W)	Altitude (m a.s.l.)	Boulder dimension (L x W x H) (m)	Sample thickness (cm)	<sup>10</sup> Be concentration (atoms g <sup>-1</sup> )	<sup>10</sup> Be concentration 1σ (atoms g <sup>-1</sup> )	<sup>10</sup> Be age (yr) (Ww) ± 1σ int (ext) error	<sup>10</sup> Be age (yr) (Pt) ± 1σ ext. error	<sup>10</sup> Be age (yr) (Kp) ± 1σ ext. error	Context
Fildes1	s08808	62° 10' 2.34" S 58° 57' 45.97" W	43	2.65 x 2.10 x 1.98	1.4	37.1 x 10 <sup>3</sup>	1.9 x 10 <sup>3</sup>	6630 ± 340 (520)	6640 ± 350	6830 ± 660	Erratic (retreat)
Fildes2	s08809	62° 08' 43.65" S 58° 55' 35.96" W	40	0.87 x 0.71 x 0.35	1.2	36.4 x 10 <sup>3</sup>	2.0 x 10 <sup>3</sup>	6510 ± 350 (520)	6530 ± 370	6710 ± 660	Erratic (retreat)
Fildes3	s08810	62° 08' 43.97" S 58° 55' 37.18" W	39	0.74 x 0.62 x 0.26	1.2	36.9 x 10 <sup>3</sup>	2.2 x 10 <sup>3</sup>	6610 ± 390 (550)	6630 ± 400	6810 ± 680	Erratic (retreat)

**Table 1.** Cosmogenic surface exposure dating results. Exposure ages were calculated using version 3 of the online calculator described by Balco et al., (2008). Following standard procedures for the Antarctic Peninsula region, we used the Antarctic (“Ant”) pressure flag and the scaling scheme “LSDn” (Stone, 2000; Lifton et al., 2014). We employed the scaling model “LSDn” because it derives from an improved understanding of the atmospheric particle flux and neutron spectrum (Lifton et al., 2014), as compared with other scaling schemes using empirical fits to proxy data. A rock density of 2.7 g cm<sup>-3</sup> was assumed for the samples. Exposure ages are derived using the default <sup>10</sup>Be production rate calibration dataset (Borchers et al., 2016) in the v. 3 of the online calculator of Balco et al (2008). However, for comparison, <sup>10</sup>Be ages were calculated using the mid-latitude southern hemisphere New Zealand (Putnam et al., 2010) and Patagonian calibration datasets (Kaplan et al., 2011) since these are proximal to Antarctica. No snow cover or erosion rate correction was applied. Internal and external uncertainties (Balco et al., 2008) are reported and we use external uncertainties to compare new exposure ages with calibrated AMS radiocarbon ages collected in this study. In the online calculator, we assumed zero erosion, used the AMS standard flag 07KNSTD, an "ant" elevation pressure flag and a density of 2.7 g cm<sup>-2</sup>. A shielding correction of 0.9999 was applied to all samples as the value is the same for all. Calculations of <sup>10</sup>Be exposure ages using the global (Ww) “LSDn” scaling (Lifton et al., 2014), the New Zealand-Macaulay (Pt) (Putnam et al., 2010), and the Patagonia calibration dataset (KP) (Kaplan et al., 2011) are shown. int: Internal error; ext: External error; ages have been rounded to the nearest 10 years. Lm production rate scaling from the online calculator (Balco et al., 2008) would give ages that are around 2% older, and thus within the uncertainties on the ages we report using LSDn. For the batch of three samples, the laboratory blank used to correct <sup>10</sup>Be concentrations was 2.06 x10<sup>-15</sup> ± 3.71x10<sup>-16</sup>, amounting to a maximum background correction of less than 6% of the total <sup>10</sup>Be atoms measured.

ID No.	Location Sample ID]	Lab. ID	Latitude (°S)	Longitude (°W)	Alt. (m a.s.l.)	Material dated	$\delta^{13}\text{C}$ (‰)	AMS $^{14}\text{C}$ age ( $^{14}\text{C}$ a)	Calibrated ages (a cal BP)					
									Max.–Min.	1 $\sigma$ %	Max.–Min.	2 $\sigma$ %	Median	Mean $\pm$ 2 $\sigma$
1	ART-03A-2015	Beta-441409	62° 10' 46.20"	58° 53' 58.34"	72	Terrestrial moss	- 22.7	1380 $\pm$ 30	1300–1260	37.9	1310–1170	95.4	1250	1240 $\pm$ 80
2	ART-03B-2015	Beta-441410	62° 10' 46.20"	58° 53' 58.34"	72	Terrestrial moss	- 20.5	1380 $\pm$ 30	1300–1260	37.9	1310–1170	95.4	1250	1240 $\pm$ 80
3	ART-01A-2015	Beta-441405	62° 10' 46.20"	58° 53' 58.34"	72	Terrestrial moss	- 22.6	1980 $\pm$ 30	1890–1820	62.6	1930–1740	95.4	1850	1850 $\pm$ 100
4	ART-01B-2015	Beta-441406	62° 10' 46.20"	58° 53' 58.34"	72	Terrestrial moss	- 22.4	1950 $\pm$ 30	1870–1850	14.8	1890–1730	95.4	1810	1810 $\pm$ 100
5	ART-01A-2015 bis	Beta-441407	62° 10' 46.20"	58° 53' 58.34"	72	Terrestrial moss	- 19.6	1920 $\pm$ 30	1890–1820	58.0	1930–1740	95.4	1840	1830 $\pm$ 100
6	ART-01B-2015 bis	Beta-441408	62° 10' 46.20"	58° 53' 58.34"	72	Terrestrial moss	- 16.8	1940 $\pm$ 30	2000–1920	60.0	2010–1870	95.4	1950	1940 $\pm$ 80
7	ART-08-ar-A-2015	Beta-441411	62° 10' 54.52"	58° 51' 43.94"	16	Terrestrial moss	- 17.9	2030 $\pm$ 30	2010–1920	68.3	2050–2030	1.8	1960	1960 $\pm$ 80
8	ART-08-ar-B-2015	Beta-441412	62° 10' 54.52"	58° 51' 43.94"	16	Terrestrial moss	- 24.5	2050 $\pm$ 30	1820–1720	68.3	1840–1690	94.4	1770	1760 $\pm$ 80
9	ART-08-ab-A-2015	Beta-441413	62° 10' 54.52"	58° 51' 43.94"	16	Terrestrial moss	- 20.9	1870 $\pm$ 30	1830–1740	68.3	1880–1700	95.4	1790	1790 $\pm$ 80
10	ART-08-ab-B-2015	Beta-441414	62° 10' 54.52"	58° 51' 43.94"	16	Terrestrial moss	- 28.6	1900 $\pm$ 30	2010–1920	68.3	2050–2030	1.8	1960	1960 $\pm$ 80
11	ART-02-1A	Beta-441415	62° 10' 56.07"	58° 53' 22.15"	7	Terrestrial moss	- 20.5	2050 $\pm$ 30	2000–1960	24.0	2000–1830	95.4	1920	1930 $\pm$ 80

12	ART-02-2A	Beta-441417	62° 10' 56.07"	58° 53' 22.15"	7	Terrestrial moss	- 21.9	2010 ± 30	1930–1830	68.3	2000–1950	12.6	1890	1890 ± 100
									–	–	1940–1820	81.6		
									–	–	1770–1750	1.3		
13	ART-02-2B	Beta-441418	62° 10' 56.07"	58° 53' 22.15"	7	Terrestrial moss	- 23.7	1980 ± 30	1930–1830	68.3	2000–1750	95.4	1890	1890 ± 100
14	ART-07A-2015	ETH-69688	62° 10' 51.73"	58° 51' 39.45"	18	<i>Laternula</i> sp.shell	-0.9	7045 ± 85	6790–6480	68.3	6940–6330	95.4	6630	6640 ± 300
15	ART-07B-2015	ETH-69689	62° 10' 51.73"	58° 51' 39.45"	18	<i>Laternula</i> sp.shell	4.0	6590 ± 70	6270–6000	68.3	6400–5880	95.4	6140	6140 ± 260
16	ART-07C-2015	ETH-69690	62° 10' 51.73"	58° 51' 39.45"	18	<i>Laternula</i> sp.shell	0.0	6730 ± 75	6440–6160	68.3	6570–6000	95.4	6300	6290 ± 280
17	ART-09A-2015	ETH-69691	62° 10' 50.59"	58° 51' 40.66"	27	<i>Laternula</i> sp.shell	-2.2	6265 ± 75	5920–5640	68.3	6090–5530	95.4	5790	5790 ± 280
18	ART-09B-2015	ETH-69692	62° 10' 50.59"	58° 51' 40.66"	27	<i>Laternula</i> sp.shell	3.4	5895 ± 70	5550–5280	68.3	5660–5080	95.4	5400	5390 ± 280
19	ART-10A-2015	ETH-69693	62° 10' 49.04"	58° 51' 34.88"	28	<i>Laternula</i> sp.shell	0.9	6235 ± 70	5890–5620	68.3	6020–5480	95.4	5760	5760 ± 260
20	ART-10B-2015	ETH-69694	62° 10' 49.04"	58° 51' 34.88"	28	<i>Laternula</i> sp.shell	-1.2	5925 ± 70	5560–5300	68.3	5700–5130	95.4	5430	5420 ± 260
21	ART-10C-2015	ETH-69695	62° 10' 49.04"	58° 51' 34.88"	28	<i>Laternula</i> sp.shell	-2.8	6520 ± 75	6200–5930	68.3	6330–5780	95.4	6070	6070 ± 260
22	ART-04A-2015	ETH-69684	62° 10' 23.69"	58° 54' 59.09"	42	<i>Laternula</i> sp.shell	-1.4	>44300 ± 4430	45510–45040	68.3	45730–44830	95.4	45280	45280 ± 460
23	ART-04B-2015	ETH-69685	62° 10' 23.69"	58° 54' 59.09"	42	<i>Laternula</i> sp.shell	2.8	>43700 ± 4370	44970–44580	68.3	45210–44430	95.4	44790	44800 ± 400
24	ART-05A-2015	ETH-69686	62° 10' 24.55"	58° 54' 59.05"	42	<i>Laternula</i> sp.shell	2.8	>48500 ± 4850	50090–49170	68.3	50800–48710	95.4	49690	49720 ± 1120
25	ART-05B-2015	ETH-69687	62° 10' 24.55"	58° 54' 59.05"	42	<i>Laternula</i> sp.shell	-	>51700 ± 5170	51900–51490	68.3	52120–51300	95.4	51700	51700 ± 400
26	K1-M [0–0.5 cm]	Beta-501582	62° 11' 36.55"	58° 57' 59.93"	15	<i>Drepanocladus</i> l. sp.	- 27.4	-100 ± 30	-6 – -7	68.3	-5 – -7	95.4	-6	-6 ± 1
27	K2-M [4.5-5 cm]	Beta-516809	62° 11' 36.55"	58° 57' 59.93"	15	<i>Drepanocladus</i> l. sp.	- 29.3	-130 ± 30	-6 – -7	68.3	-6 – -7	95.4	-7	-7 ± 1

28	K3-BS [5-6 cm]*	Beta-516810	62° 11' 36.55"	58° 57' 59.93"	15	Bulk sediment	- 18.4	1350 ± 30*	650–460	68.3	750–370	95.4	560	560 ± 180
29	K4-BS [9–10 cm]*	Beta-504801	62° 11' 36.55"	58° 57' 59.93"	15	Bulk sediment	- 23.9	1600 ± 30*	910–700	68.3	1010–600	95.4	800	800 ± 200
30	K5-BS [20–21 cm]*	Beta-497408	62° 11' 36.55"	58° 57' 59.93"	15	Bulk sediment	- 32.6	9380 ± 30*	10030–9790	68.3	10130–9660	95.4	9910	9900 ± 240
31	K6-BS [36.5–37 cm]	Beta-501584	62° 11' 36.55"	58° 57' 59.93"	15	Bulk sediment	- 18.0	1860 ± 30	1820–1770	27.8	1830–1690	90.6	1750	1750 ± 100
32	K7-BS [45.5–46 cm]	Beta-501585	62° 11' 36.55"	58° 57' 59.93"	15	Bulk sediment	- 17.7	3350 ± 30	1750–1700 3570–3480	40.5 68.3	1650–1620 3680–3670	4.9 1.0	3530	3530 ± 100
33	K8-BS [73.5–74.5 cm]	Beta-498331	62° 11' 36.55"	58° 57' 59.93"	15	Bulk sediment	- 18.0	6410 ± 30	– 7420–7390 7330–7250	– 9.1 29.2	3640–3450 7430–7240 7210–7170	94.5 85.5 10.0	7300	7300 ± 120

**Table 2.** New radiocarbon age data from the eastern Shetland I Moraine and the Kiteschsee Lake record on the Fildes Peninsula. Calibrated ages have been rounded to the nearest 10 years. Dating replicates were used to assess dating and interpretation accuracy. Elevation data (Alt. m a.s.l. – Altitude in metres above present sea level) were taken with a Garmin handheld GPS map 62stc, precision ±2 m (latitude, longitude) and altitudinal precision of ±5–10 m. Calibrated and modelled radiocarbon ages have been rounded to the nearest 10 years, with two-sigma calibrated ages or 95% confidence age-depth model age ranges shown and quoted in the text.



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