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**Advances in understanding subglacial meltwater drainage
from past ice sheets**

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1 Advances in understanding subglacial meltwater drainage from past ice sheets

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13

14 Abstract

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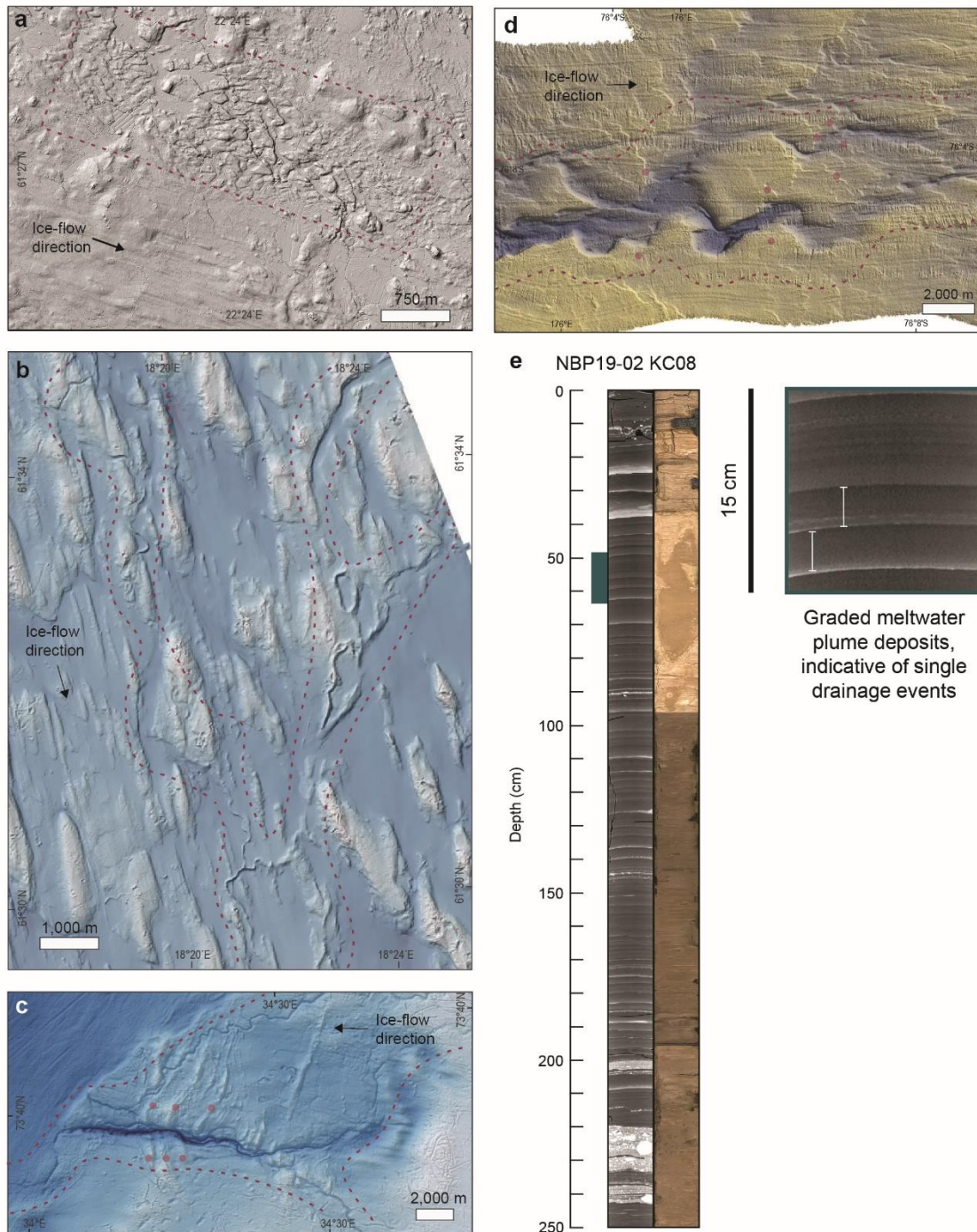
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25 1. Introduction

26 Liquid water beneath ice sheets influences ice-flow organization and velocity (Bell et al., 2007;
27 Kyrke-Smith et al., 2014; Larter et al., 2019), subglacial sediment rheology and transport (Damsgaard et
28 al., 2020; Minchew & Meyer, 2020), grounding-line behavior (Horgan et al., 2013; Fried et al., 2015), and
29 ice-shelf mass balance (Le Brocq et al., 2013; Alley et al., 2016). Ice-sheet response, however, is
30 contingent on subglacial water supply and drainage organization (Röthlisberger, 1972; Walder, 1986;
31 Schoof, 2010). While some components of subglacial hydrological systems are relatively stable (i.e.,
32 fixed), such as large subglacial lakes beneath the East Antarctic Ice Sheet (Kapitsa et al., 1996) and
33 incised bedrock channels (Kirkham et al., 2020), other reservoirs and drainage pathways are more
34 transient and evolve through time and space with nonlinear and spatially heterogeneous impacts on ice-
35 sheet behavior (Schroeder et al., 2013; Andrews et al., 2014; Hoffman et al., 2016; Siegfried et al., 2016;
36 Rada & Schoof, 2018). Beyond the grounding line, contemporary sediment plumes emanating from
37 marine-terminating outlet glaciers of the Greenland Ice Sheet observed via satellite imagery (Fried et al.,
38 2015; Schild et al., 2016) and surficial expressions of channelization beneath Antarctic ice shelves (Le
39 Brocq et al., 2013; Alley et al., 2016) indicate active subglacial hydrological systems upstream.

40 Major advances in observing contemporary ice-sheet hydrology, such as radar specularly
41 (Schroeder et al., 2013) and repeat satellite measurements (Fricker et al., 2015), reveal spatiotemporal
42 evolution of basal water transmission on sub-decadal scales. Yet, limited long-term (decadal to millennial)
43 observations impede holistic perspectives on the modes and magnitudes of water drainage beneath ice
44 sheets and their consequences for ice-sheet behavior. In formerly glaciated landscapes and continental
45 margins, relict subglacial water drainage is recorded by meltwater landforms and sedimentological
46 successions (Fig. 1; Kehew et al., 2012; Lee et al., 2015; Esteves et al., 2017; Greenwood et al., 2016).

47 Channels (broadly defined) incised into bedrock and sediments and positive-relief esker ridges record
 48 drainage styles and organization (Storrar et al., 2014; Zoet et al., 2018; Lewington et al., 2020) and, in
 49 some cases, associated ice-margin retreat behavior (Livingstone et al., 2020; Simkins et al., 2021).
 50 Distinct meltwater plume deposits and hydrologically sorted sediments reveal relative magnitudes and
 51 frequency of water drainage into the ocean, precursory, synchronous, or resulting glacial environment
 52 changes, and geochemical signatures of sediment and water provenance (Witus et al., 2014; O'Regan et
 53 al., 2021; Lepp et al., 2022). Such empirical observations based on the geological record have long
 54 guided and continue to inform, and challenge, glaciological theory (e.g., Walder & Hallet 1979; Boulton et
 55 al. 2009; Hewitt, 2011).



56 **Fig. 1.** (a) Murtoo pathway within glacially streamlined terrain in central Finland (Mäkinen et al. 2017; Ojala et al. 2019). Data:
 57 LiDAR-based DEM from the National Land Survey of Finland. (b) Meltwater channels and eskers drape and incise drumlins in the
 58 Bothnian Sea (Greenwood et al. 2017). Data: MBES-based DEM from the Swedish Maritime Administration. (c) Meltwater channel
 59 incised retreat moraines (red dots) on Thor Iversenbanken in the Central Barents Sea (Esteves et al., 2017). Data: MAREANO
 60 MBES-based bathymetry from the Norwegian Mapping Authority. (d) Meltwater corridor in which channels cross-cut grounding zone
 61 wedges (red dots) in the Ross Sea, Antarctica (Simkins et al. 2021). Data: MBES-based DEM from cruise NBP15-02, available
 62 through the United States Antarctic Program Data Center. In (a)-(d), red dashed lines outline the encompassing areas of meltwater
 63 landforms. (e) Computed tomography (CT) scan and photograph of the upper 250 cm of sediment core NBP19-02 KC-08, collected
 64 in the Amundsen Sea, records meltwater plume events that emanated from the Thwaites Glacier grounding line (Lepp et al. 2022).
 65

66 2. Recent advances from paleo-ice sheets

67 Over the past decade, advances in geophysical methods, growing data accessibility, and
 68 sedimentological studies from deglaciated terrains push the boundaries of subglacial hydrology
 69 understanding and challenge concepts of the spatiotemporal evolution of water drainage beneath ice
 70 sheets. Near complete coverage of light detection and ranging (LiDAR) and satellite photogrammetry
 71 elevation data across terrestrial landscapes formerly glaciated by the European and North American ice
 72 sheets, gives unprecedented views of paleo-subglacial meltwater landforms that range in relief from 10^{-1}
 73 to 10^2 of meters and lengths of 10^1 to 10^5 meters, allowing holistic ice-sheet scale assessment of controls
 74 on drainage. While nowhere near as complete, increasing coverage and quality of bathymetry data from
 75 deglaciated continental shelves provides perspectives on water flow beneath marine-based ice sheets
 76 and implications for ice-sheet behavior at fine scales previously unseen. These offshore advances via
 77 multibeam echo sounding (MBES) surveys are facilitated by national hydrographic programs such as
 78 MAREANO (e.g., Esteves et al., 2017), marine geological repositories such as the Marine Geoscience
 79 Data System, and researcher-led surveying and compilations (e.g., Greenwood et al., 2021). Additionally,
 80 3-D seismic survey grids, albeit sparse but increasing in spatial coverage due to industry-academic
 81 relations, are unique datasets to assess temporal evolution of drainage pathways and internal
 82 architecture of meltwater landforms (e.g., Kirkham et al., 2021, 2022). These advances in terrain data
 83 acquisition and availability have not only permitted ice-sheet scale documentation of small-scale (meter to
 84 sub-meter) and intricate meltwater landforms, demonstrating their variability as well as near ubiquity, but
 85 have also uncovered both new types of landforms and little recognized meltwater landform assemblages,
 86 stimulating new hypotheses for meltwater landform genesis and understanding of the coupling to ice-flow
 87 and ice-margin behavior (Storrar et al., 2014; Ojala et al., 2019; Kirkham et al., 2020).

88 Challenging the binary categorization of subglacial drainage through either “channelized” or
 89 “distributed” pathways, meltwater corridors found in the Northern Hemisphere (e.g., Peterson et al., 2018;
 90 Lewington et al., 2020), and on the Antarctic seafloor (Simkins et al., 2021) represent broad subglacial
 91 drainage pathways of meltwater landform assemblages that span 10^1 - 10^2 kilometers in length. These
 92 corridors indicate co-existence of drainage styles, varying genetic erosional and depositional processes,
 93 and waxing and waning of drainage magnitudes in time and space. Additionally, newly observed
 94 landforms in Scandinavia, termed murtoos, potentially bridge the long-standing gap in recognizing
 95 geomorphic evidence for distributed subglacial drainage (Mäkinen et al. 2017; Ojala et al., 2019). These
 96 low-relief triangular subglacial landforms oriented with their apex in the ice-flow direction (Fig. 1a) often
 97 occur with other meltwater landforms such as channels and eskers, and sediments which have
 98 undergone hydraulic sorting, ductile deformation, and liquefaction (Becher & Johnson, 2021). Collectively,
 99 murtoo presence suggests efficient transitional drainage between channelized and distributed under high
 100 pressure conditions, potentially in response to transient linked cavity-type drainage systems (Ojala et al.,
 101 2022) similar to those beneath the Greenland Ice Sheet (Hoffman et al., 2016). Both corridors and
 102 murtoos point to variable modes of drainage that co-exist or evolve in time and space including “efficient”
 103 and “inefficient” components, thus questioning the validity of assuming or parameterizing singular modes
 104 of subglacial water drainage.

105 A long-standing challenge in glacial geomorphology has been how to interpret the temporal
106 significance of meltwater landforms: the time required, and the stability of discharge required, for both
107 landform and whole drainage pathway formation. Meltwater landform relations to other subglacial and ice-
108 marginal landforms provide insights in this regard (e.g., [Simkins et al., 2017](#); [Greenwood et al., 2017](#);
109 [Ojala et al., 2019](#); [Livingstone et al., 2020](#)). For example, drumlins and mega-scale glacial lineations
110 incised by channels and draped by eskers in the Bothnian Sea indicate a geomorphic switch from active
111 bedform shaping to channelized water drainage overprinting stable bedforms, shortly before deglaciation
112 ([Fig. 1b](#); [Greenwood et al., 2017](#)). Here, interlinking channels and eskers of comparable sizes within a
113 coherent drainage path highlight the transitory dominance of erosion and deposition in the subglacial
114 environment. Episodic esker segment (“bead”) deposition has long been inferred from the terrestrial
115 landform-sediment record ([De Geer, 1897](#); [Banerjee and McDonald, 1975](#); [Mäkinen, 2003](#)). [Livingstone
116 et al. \(2020\)](#) demonstrate a tight relationship between esker beads and De Geer moraines in central
117 Nunavut and infer time-transgressive landform building, yet relatively fixed in space, by drainage
118 pathways to the ice margin. Similarly, meltwater channel incision through retreat moraines in the Barents
119 Sea ([Fig. 1c](#); [Esteves et al., 2017](#)) and variable incision of or draping by retreat moraines in the western
120 Ross Sea ([Simkins et al. 2017](#)) indicate the relative persistence of channelized drainage during active ice-
121 margin retreat. A corridor of over 80 meltwater channels on the Antarctic continental shelf ([Fig. 1d](#);
122 [Simkins et al., 2021](#)) had prolonged impacts on grounding-line behavior as larger magnitude grounding-
123 line retreat events and grounding zone wedge deposition occurred while the channels within the corridor
124 were active, compared to smaller retreat events and moraine deposition when the channels were inactive.
125 While the mechanism for this relationship remains unknown, it possibly results from hydrological controls
126 on sediment rheology and mobility that influence building of ice-marginal landforms that may or may not
127 reduce effective water depths enough to counterbalance grounding line buoyancy. Such observations of
128 meltwater - grounding line landform associations and the potential to document these over large tracts of
129 paleo-ice sheet beds offers new possibilities for constraining the time component of the meltwater
130 landform record, as well as quantifying sediment loads and, for example, seasonal deposition of individual
131 esker beads.

132 Complementary to geomorphological studies, sediment records from deglaciated continental
133 shelves and proglacial lake basins elucidate the temporal persistence of subglacial and grounding line
134 water discharge and associated changes in ice-sheet configuration and behavior (e.g., [Rüther et al.,
135 2012](#); [Lee et al., 2015](#); [Avery et al., 2021](#); [O'Regan et al., 2021](#); [Lepp et al., 2022](#)). Meltwater plume
136 deposits offshore of Thwaites Glacier, Antarctica and Ryder Glacier, Greenland are a common feature
137 associated with (or precursor to) glacier retreat and ice-shelf break up events, indicated by the millimeter-
138 scale stratigraphy resolved by computed tomography (CT) scans and by grain-scale sedimentology ([Fig.
139 1e](#); [O'Regan et al., 2021](#); [Lepp et al., 2022](#)). Downcore stratigraphy and trace elemental ratios in cores
140 that sample meltwater plume deposits reveal differences in relative magnitudes and frequencies of
141 subglacial drainage into the ocean offshore of western and eastern Thwaites Glacier, and suggest greater
142 magnitudes of sediment-laden water were delivered to the ocean in recent centuries compared to the
143 past several thousand years ([Lepp et al., 2022](#)). In the Baltic Sea basin, where proglacial varved
144 sediments have long been used to document the pattern and pace of Fennoscandian ice-margin retreat,
145 [Avery et al. \(2021\)](#) find multi-decadal cycles of enhanced meltwater discharge through a 725-year varve
146 series, around 15,000 years ago. In these paleo cases, and particularly where the former ice sheet bed is
147 now exposed, there is great potential for examining links between the temporal information archived in
148 the distal sedimentological record of meltwater events and longevity of discharge, and the high-resolution
149 geomorphology of the hydrological system responsible.

150 A recent body of work examining physical processes of and conditions for subglacial fluvial
151 erosion, deposition and sediment mobility is an important step forward ([Beaud et al., 2016, 2018](#);
152 [Damsgaard et al., 2017](#); [Hewitt and Creyts, 2019](#); [Kirkham et al., 2022](#); [Vérité et al., 2022](#); [Stevens et al.,](#)

153 2022). These studies build towards an integrated or continuum view of meltwater organization, depending
154 on water supply and sources, basal conditions and substrate properties. Importantly, they make advances
155 towards knowledge of where, over what timescales, and with what meltwater discharge regimes,
156 sediments are mobilized and landforms may form, opening up the vast landform record to much more
157 effective and accurate use as a document of coupled meltwater – ice flow – ice margin behavior in paleo-
158 ice sheets over seasonal-to-millennial timescales.

159

160 3. Looking forward

161 To push the field of subglacial hydrology forward using the landform and sediment records of
162 deglaciated regions, we need: increased geophysical data coverage in regions proximal to contemporary
163 ice-sheet margins; coupled remote-sensing and field-based observations in terrestrial landscapes;
164 reporting of quantitative sedimentologic and morphometric data; and community building to work across
165 disciplinary bounds and study-area silos. Of promise in narrowing knowledge gaps are emerging
166 themes of research on deeper groundwater interactions with the ice-bed interface (Gustafson et al., 2022)
167 and its implications for landform genesis (Boulton et al., 2009; Hermanowski and Piotrowski, 2019),
168 understanding subglacial lake and ice-sheet surface connections to subglacial drainage systems
169 (Greenwood et al., 2016; Simkins et al., 2017), assessment of the role of local (10^0 - 10^1 m relief) variability
170 in bed conditions on drainage organization (Simkins et al., 2021), and the continued pursuit of
171 constraining time for meltwater landform construction and evolution; each of these pursuits will benefit
172 from more seamless integration of theory, numerical models, and coupled geomorphological and
173 chronological studies (Kirkham et al., 2022; Stevens et al., 2022). Additionally, higher-resolution
174 topographic data from the surface of Mars offers opportunities to dig deeper into the surficial expressions
175 of meltwater landforms (e.g., Butcher et al., 2020), whereby comparison with Earth's meltwater landforms
176 may offer new insights into key processes controlling their genesis. Here on Earth and on Mars, we need
177 to be mindful of what we are not seeing, such as evidence for distributed, transient, and hard-bed
178 systems that might not leave their mark, in geomorphological and sedimentological records of subglacial
179 hydrological systems. Additionally, when planning new research projects, those of us working on records
180 of paleo-ice sheets and those studying contemporary ice sheets should draw on literature from the two
181 respective fields to identify key gaps in understanding of subglacial hydrological forms and processes that
182 will aid in assessing the future of the Greenland and Antarctic ice sheets.

183

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189

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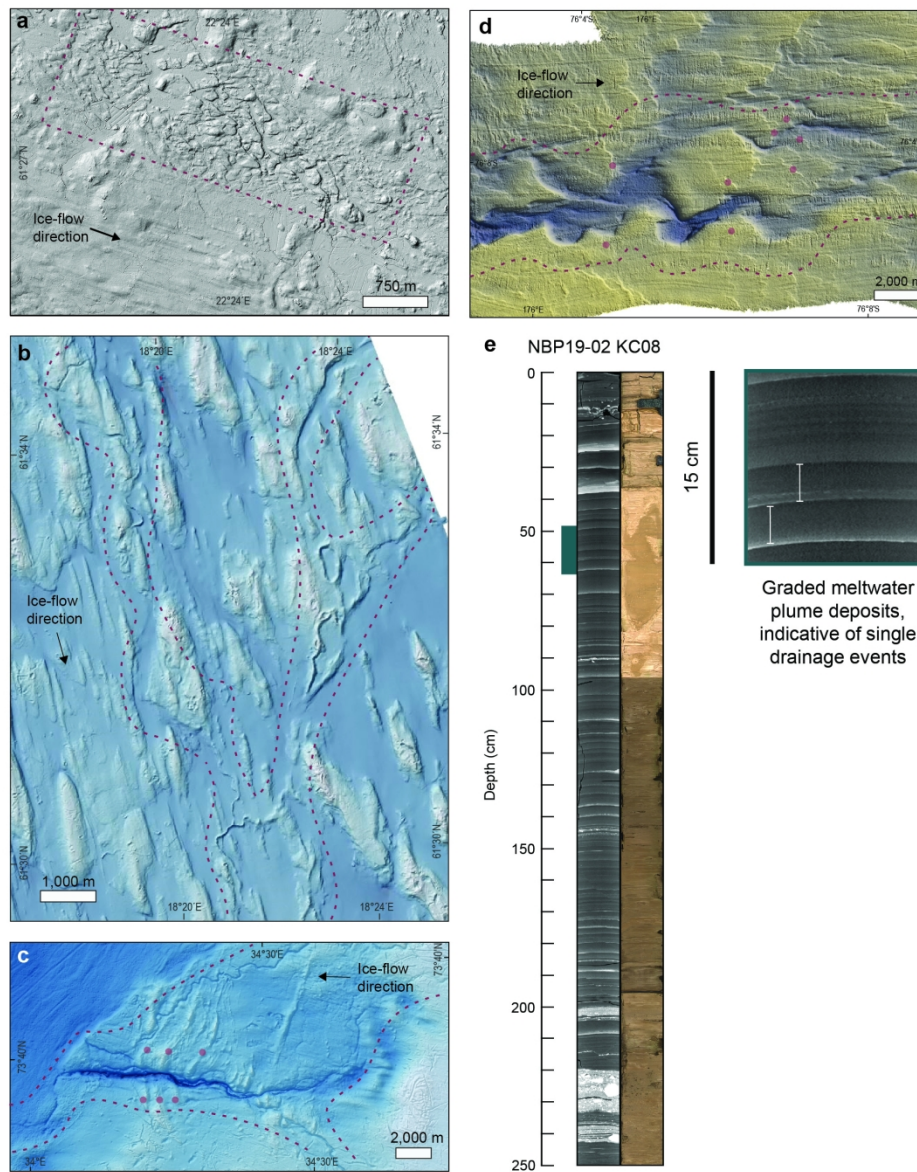


Fig. 1. (a) Murtoo pathway within glacially streamlined terrain in central Finland (Mäkinen et al. 2017; Ojala et al. 2019). Data: LiDAR-based DEM from the National Land Survey of Finland. (b) Meltwater channels and eskers drape and incise drumlins in the Bothnian Sea (Greenwood et al. 2017). Data: MBES-based DEM from the Swedish Maritime Administration. (c) Meltwater channel incised retreat moraine (red dots) on Thor Iversenbanken in the Central Barents Sea (Esteves et al., 2017). Data: MAREANO MBES-based bathymetry from the Norwegian Mapping Authority. (d) Meltwater corridor in which channels cross-cut grounding zone wedges (red dots) in the Ross Sea, Antarctica (Simkins et al. 2021). Data: MBES-based DEM from cruise NBP15-02, available through the United States Antarctic Program Data Center. In (a)-(d), red dashed lines outline the encompassing areas of meltwater landforms. (e) Computed tomography (CT) scan and photograph of the upper 250 cm of sediment core NBP19-02 KC-08, collected in the Amundsen Sea, records meltwater plume events that emanated from the Thwaites Glacier grounding line (Lepp et al. 2022).

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