Coversheet

Dissolution methods for strontium isotope stratigraphy: Guidelines for the use of argillaceous, dolomitic limestone

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Dissolution methods for strontium isotope stratigraphy: Guidelines for the use of argillaceous, dolomitic limestone

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Abstract

In recent years, various sample preparation methods have been developed to extract a primary Sr isotope signal from carbonate rocks. However, there has been little consensus around the best method due to natural variability in sample purity and mineralogy. For this study, we conducted systematic leaching experiments, focussing mainly on generally less favoured argillaceous and dolomitic limestones, using samples from the Mesoproterozoic Gaoyuzhuang Formation. Learning from previous studies, a multistep leaching approach was used to explore strontium isotope leaching systematics and extract near-primary isotopic ratios. For argillaceous and dolomitic limestone, the first 10%-30% dilute acetic acid leach, following an ammonium acetate (NH_4Ac) prewash, yielded the lowest, demonstrably least altered seawater $^{87}Sr/^{86}Sr$ ratios. Subsequent dissolution steps exhibited significantly elevated $^{87}Sr/^{86}Sr$, Rb/Sr, Al/Ca, and Mg/Ca ratios, indicating greater contributions from aluminosilicates and dolomite in the leachates. Lanthanide rare earth element and yttrium (REY) concentrations were also examined for each leaching step. In most cases, seawater-like REY patterns with the highest Y/Ho ratios (mostly > 36) occur in the first leaching step after ammonium acetate prewash, which is consistent with the Sr isotope study. However, one organic-rich sample (TOC_{total} ~1.3%) with a near-seawater Sr isotope value exhibits a non-seawater-like REY pattern, possibly indicating that early diagenetic exchange during organic remineralization may influence the REY pattern

of carbonate rocks without necessarily changing Sr isotopes. We applied previously proposed leaching cut-offs (preleach 30% ~ 40%, 60% ~ 70%), alongside the cut-off proposed from this study, for samples with various Mg/Ca and carbonate purities. We found that for the same argillaceous and dolomitic carbonate rocks, applying higher leaching cut-offs might cause sizeable error compared with targeting the first 10% ~ 30% directly after NH_4Ac prewash. However, no significant differences were evident for high purity, and low Mg/Ca limestones, highlighting the importance of matching different sample types to the most appropriate dissolution method. Here we demonstrate that argillaceous, even partially dolomitized limestones could help fill gaps in the Precambrian seawater Sr isotope curve, especially where higher purity limestone successions are not readily available.

Key Words

Dissolution methods; strontium isotopes; argillaceous limestone; dolomitic limestone

1. Introduction

The secular trend of seawater strontium isotope ratio ($^{87}Sr/^{86}Sr$) reflects variations in the relative contributions of continental versus mantle reservoirs to ocean composition (Spooner, 1976; Veizer, 1989). The ocean residence time of Sr is much longer than the ocean circulation time (~10^6 vs. ~10^3 years, respectively), and so strontium isotopes are believed to be homogeneously distributed in seawater at any given time (Broecker and Peng, 1983; Elderfield, 1986; Hodell et al., 1990). As a consequence, strontium isotope stratigraphy (SIS) has been widely used for chemostratigraphic correlation (Burke et al., 1982; Elderfield, 1986; McArthur, 1994; Veizer et al., 1999); dating by comparison to standard reference curves (McArthur et al., 2012, 2020); tracing diagenetic processes and depositional environment (Kuznetsov et al., 2010; Stüeken et al., 2017); and testing hypotheses of tectonic, biological and climatic evolution on geological timescales (e.g., Halverson, 2007; Shields, 2007; Hawkesworth et al., 2016; Cawood et al., 2018).

SIS studies must rely on diagenetically well-preserved chemical precipitates and well-honed dissolution methods. Diagenetic processes tend to increase $^{87}Sr/^{86}Sr$ values when interstitial fluids are influenced by evolved K-bearing silicates (e.g., Shields and Veizer, 2002; Fairchild et al., 2018), or decrease Sr isotope composition when diagenetic fluids are influenced by mafic components, hydrothermal fluids or pressure solution of older, underlying carbonate rocks (e.g., Miller et al., 2008; Brand et al., 2010; Satkoski et al., 2017; Cui et al., 2020). Unfortunately, the well preserved, low-Mg calcite fossils, widely used in Phanerozoic SIS studies, are not available in Precambrian rocks, and so fine-grained carbonate components (e.g., diagenetic calcite microspar cement; Zhou et al., 2020), bulk carbonate rocks (e.g., micrite; Bailey et al., 2000) or non-carbonate rocks such as barite (e.g., McCulloch, 1994; Satkoski et al.,

2016; Roerdink et al., 2022), gypsum or anhydrite (e.g., Kah et al., 2001) and francolite (Li et al., 2011) have all been used instead for this purpose. Apart from diagenetic alteration, the leaching of detrital aluminosilicate phases during sample preparation can also introduce unintended Sr contamination, which is either released by ion exchange during the initial leaching stage or aluminosilicate dissolution during the later leaching process (McArthur, 1994; Bailey et al., 2000; Bellefroid et al., 2018). Clay contamination during sample preparation normally would lead to Sr isotopes much higher than expected (Bailey et al., 2000), may result in overcorrection for radioactive Rb decay (Shields and Veizer, 2002), and can mask the original diagenetic trend in a suite of samples (Bellefroid et al., 2018). Therefore, when conducting SIS, it is crucial that Sr is isolated from targeted and least altered minerals without contamination from extraneous phases, which requires appropriate dissolution methods.

The dissolution methods used in SIS studies can be divided into three main types: single-step bulk leaching methods, two-step sequential leaching methods and multiple-step sequential leaching methods (see recent review by Chen et al., 2022 and references therein). The most commonly used and effective sequential leaching method uses dilute acetic acid to target a certain proportion of pure carbonate, following an initial pre-leach with ammonium acetate or acetic acid to remove exchangeable Rb and Sr, while aiming to leave at least 10% carbonate undissolved (Bailey et al., 2000). This method has been developed further and is now widely used to extract primary carbonate signals from different types of carbonate rocks, using various pre-leaching cut-offs. For example, Bailey et al. (2000), confirmed subsequently by Li et al., (2011), suggested dissolving 30% ~ 40% of the carbonate portion of a sample, before targeting the following 30% ~ 40% for isotopic analysis.

Liu et al., (2013) suggested to pre-leach up to 70% ~ 80% of a dolostone sample, targeting the next 10% ~ 20% for analysis. Finally, Li et al., (2020) suggested an acidic pre-leach to dissolve ~60% of the carbonate, before targeting the next ~20% for both limestone and dolostone samples, but with emphasis on sample purity > 75% and > 90% for limestone and dolostone, respectively.

From the above, it is evident that few of any studies have sought to improve Sr isotope leaching methods for low purity and/or partially dolomitised samples, despite the lack of high purity limestones and large data gaps in Precambrian successions (Shields and Veizer, 2002; Chen et al., 2022). Moreover, without a clear definition/classification system for sample purity and sample types, the application thresholds between different methods remain ill-defined, with little consensus as to the most appropriate method. Therefore, it is imperative to conduct a systematic methodological study on the abundant but less favoured argillaceous (often referred to as ‘dirty’ or ‘muddy’) and/or dolomitic limestones, and also test different leaching cut-offs / methods for samples using clear purity and Mg/Ca classification.

The rare earth element (REE) plus yttrium (REY) compositions of marine carbonate rocks could reveal depositional environment, redox conditions and diagenetic alteration (e.g., Satkoski et al., 2017; Verdel et al., 2018), thus might help us to determine if the Sr isotope data record seawater signal. The typical seawater REY profile shows progressive enrichment in heavier REE (James et al., 1995), while REY carbonate systematics are relatively resistant to diagenetic exchange compared with Sr isotopes due to the high partition coefficients of REY between calcite and seawater and generally low REY concentrations in diagenetic fluids (Zhong and Mucci, 1995; Webb and Kamber, 2000). However, carbonate REY components are very sensitive to detrital contamination due to the much higher REY contents and

distinctly different shale - normalized REE patterns in detrital minerals (e.g., Nothdurft et al., 2004). Such a high sensitivity for clay contamination is also expected for Sr isotopes, and so comparing the REY pattern and Sr isotope ratios of individual leaching steps could help to validate the leaching method, especially for low purity carbonate rocks (James et al., 1995). The relatively clean proportion of the argillaceous and dolomitic carbonates during leaching procedures might be expected to demonstrate the most pristine, 'marine' REY patterns and Sr isotope values.

Based on earlier studies, such as Chilingar, (1957) and Zhou et al., (2020), we divided samples into four types according to their Mg/Ca weight ratios, which are limestone (LST, Mg/Ca < 0.025), slightly dolomitic limestone (SDL, 0.025 < Mg/Ca < 0.25), highly dolomitic limestone (HDL, 0.25 < Mg/Ca < 0.6) and dolostone (DST, Mg/Ca > 0.6). We define sample purity < 80% as argillaceous/low purity samples, and sample purity > 80% as relatively pure samples. In this study, we mainly aim to 1) conduct a systematic method study focusing mainly on argillaceous and dolomitic limestones (sample purity < 80%, 0.025 < Mg/Ca < 0.6); 2) use REY patterns to assist in the discussion of Sr isotope leaching results and explore the connections and differences between these two systems; 3) explore the fitness of different methods for different types of rock by applying previously published, commonly used leaching cut-offs to samples with various purities and Mg/Ca ratios; and 4) propose new thresholds for sample screening in future SIS studies.

2. Geological setting and sample descriptions

2.1. Geological background

The ~1.65-1.4 Ga Jixian Group consists of five formations (Gaoyuzhuang, Yangzhuang, Wumishan, Hongshuizhuang, and Tieling in ascending order), which represent near-continuous marine deposition within the Yanliao Basin, North China Craton. The ~1.6-1.55 Ga Gaoyuzhuang Formation is further divided into four lithological members in ascending order: Guandi Member (M1), Sangshu'an Member (M2), Zhangjiayu Member (M3) and Huanxiusi Member (M4). Most samples in this study were collected from the ~1.57-1.56 Ga Zhangjiayu Member (M3) (Li et al., 2010; Tian et al., 2015), through an interval that covers a negative carbon isotope excursion (Li et al., 2003; Guo et al., 2013; Zhang et al., 2018). The Zhangjiayu Member consists mainly of limestone and dolomitic limestone deposited in the shallow marine environment, with argillaceous and variably dolomitised limestone and black shale mainly in the lower part, showing a shallowing-upward cycle (e.g., Mei et al., 2005; Zhang et al., 2018). The impure, dolomitic limestones of Zhangjiayu Member provide a good opportunity for us to examine the leaching method for this type of rock. Based on a new compilation of the Precambrian seawater Sr isotope curve (Chen et al., 2022), the published and screened seawater $^{87}Sr/^{86}Sr$ data between 1.60 Ga and 1.55 Ga range from ~0.7046 to ~0.7056 (e.g., Ray et al., 2003; Kuznetsov et al., 2008; Bellefroid et al., 2018; Tan et al., 2020), which are compared with data obtained in this study.

2.2. Samples

Samples of varying purity and dolomitization were collected from the Zhangjiayu Member (M3) at Jixian (J), Pingquan (P), Gangou (G), Kuancheng (K), Huanxiusi

(HXSZ) and Sangshu'an (SSAZ) sections, and underlying Huanxiusi Member (M4, abbreviated to HXS) on the North China Craton. Samples analyzed in this study were micro-drilled to extract powder to avoid visibly secondary portions such as cross-cutting veins, late-stage void filling spar and dissolution features etc. After that, powders of each sample were dissolved in 2% HNO_3 for bulk carbonate major and trace elements before undergoing leaching steps. Stable isotopes ($\delta^{13}C$ and $\delta^{18}O$) were also analyzed to help with sample selection and diagenetic screening. Five samples (J95, G99, K46, G90, HXSZ4) from four different sections (Jixian, Gangou, Kuancheng and Huanxiusi) of Zhangjiayu Member (M3) were selected to develop a leaching method for argillaceous and dolomitic limestone. These samples have Mg/Ca ratios ranging from ~0.06 to 0.4 with carbonate contents ranging from 50% to 75%, and with few obvious signs of significant diagenetic alteration (e.g., in most cases Mn/Sr <1, $\delta^{13}C$ close to ~0, $\delta^{18}O$ at -3 ~ -5‰). The five samples were used for sequential leaching experiments alongside rock standard LS19 (a pure limestone from the Huaibei group, Mg/Ca = 0.02) from Zhou et al., (2020) with a published Sr isotope value of 0.705439 (±3); and one reference material CRM 88a (a pure dolostone, Mg/Ca = 0.6) with a published Sr isotope value of 0.71022 (Stammeier et al., 2020). The detailed bulk rock information for all samples used for this study is summarized in **Table 1**. Another twenty samples with various sample purities and Mg/Ca ratios were studied to examine the leaching methods, the bulk carbonate information of which is reported alongside Sr isotope result in **Table 4** in section 4.3.

Sample Labels	Sample types	Carbonate content (%)	Mg/Ca (g/g)	Mn/Sr (g/g)	Rb/Sr (mg/g)	K/Ca (mg/g)	Al/Ca (mg/g)	Ba/Ca (mg/g)	Fe/Ca (mg/g)	[Sr/Ca+Mg] (ppm)	$\delta^{13}C_{VPDB}$ (‰)	$\delta^{18}O_{VPDB}$ (‰)	Age (Ga)
LS19	LST	90%	0.02	0.03	0.1	0.38	0.39	0.01	0.10	988			~0.96
J95	SDL	70%	0.06	0.14	5.48	0.79	2.15	0.06	2.02	672	-0.64	-5.25	~1.56-1.57
G99	SDL	57%	0.11	0.51	36.90	4.75	5.73	0.19	17.84	876	-1.27	-4.89	~1.56-1.57
K46	SDL	75%	0.20	0.34	11.47	1.73	3.43	0.09	5.91	661	-0.59	-4.52	~1.56-1.57
G90	HDL	62%	0.40	1.31	30.67	2.01	2.80	0.09	14.12	368	-0.99	-3.49	~1.56-1.57
HXSZ4	HDL	62%	0.30	0.28	6.88	1.28	2.46	0.11	6.89	519	-0.52	-4.21	~1.56-1.57
88a	DST	98%	0.6	4.83	3.02	4.60	4.00	0.01	9.00	121			

Table 1. Bulk carbonate information for samples used for step-leaching experiments.

3. Methods

3.1. Sequential leaching procedure

3.1.1. Ten-step sequential leaching procedure

Micro-drilled sample powder was ground by hand using an agate pestle and mortar to avoid coarse flakes, and then around 100-200 mg of each sample was transferred to 10 ml centrifuge tubes to carry out the sequential leaching procedure. Modified after Bailey et al., (2000), a ten-step leaching procedure was applied to argillaceous, dolomitic limestones (J95, G99, K46, G90, HXSZ4) and CRM 88a, while a six-step leaching procedure was applied to pure limestone LS19. All samples have been prewashed with 1M ammonium acetate, followed by dilute acetic acid (0.05 M-0.5 M) dissolution. In each step, based on total carbonate content, the acid volumes were designed to dissolve ~10% carbonate for argillaceous and dolomitic limestones and CRM 88a, and ~20% for pure limestone LS19. After the acid was added, samples were ultrasonically agitated (0.5 hr-2 hr) and allowed to stand for 0.5 hr-2 hr at 25-40°C, then were centrifuged at 3600 rpm for 5mins. The supernatant was collected for Sr isotope and elemental analysis, and the residue washed with ultrapure water one time before the next leaching step. The carbonate proportion being leached out in each step was calculated by using mass of ($CaCO_3$+$MgCO_3$) in each leaching step divided by mass of ($CaCO_3$+$MgCO_3$) in the whole sample. In total, more than 95% of carbonate was dissolved from each sample. Detailed leaching protocol is summarized in **Table 2.**

Steps	LST	SDL	HDL	DST	Reaction time
S0	5 ml 1 M NH_4HAc	5 ml 1 M NH_4HAc	5 ml 1 M NH_4HAc	5 ml 1 M NH_4HAc	30 mins usb at room temperature
S1-S2	7.5 ml 0.05 M HAc	4 ml 0.2 M HAc	2.5 ml 0.5 M HAc	5 ml 0.5 M HAc	30 mins usb + 30 mins stand at room temperature
S3-S5	7.5 ml 0.05 M HAc	5 ml 0.2 M HAc	3.5 ml 0.5 M HAc	8 ml 0.5 M HAc	30 mins -1 hr usb + 30 mins -1 hr stand at room temperature
S6-S8		8 ml 0.2 M HAc	6 ml 0.5 M HAc	10 ml 0.5 M HAc	2 hr usb + 2 hr stand at 30-40℃
S9		10 ml 0.5 M HAc	10 ml 0.5 M HAc	10 ml 0.5 M HAc	Overnight on mixing roller at room temperature

Table 2. Sequential leaching protocol. 6-step leaching (design to dissolve 20% carbonate in each step) was conducted for pure limestone (LST); and 10-step leaching (designed to dissolve 10% carbonate in each step) was conducted for SDL, HDL, DST. *p*H of each reagent: 1 M NH_4HAc (~7); 0.05 M HAc (~3); 0.2 M HAc (~2.7); 0.5 M HAc (~2.5). usb: ultrasonic bath.

3.1.2. Leaching methods' comparison

Twenty samples with various sample purities and Mg/Ca from 0 to 0.6 (i.e., limestone, slightly, and highly dolomitic limestone) from Gaoyuzhuang Formation were selected to examine the suitability of various leaching cut-offs for different types of rocks (note: samples with Mg/Ca > 0.6 are not examined in this study). The following leaching methods are compared: *method 1*- prewash using ammonium acetate, then target the first 10% ~ 30% (this study); *method 2* - preleach 30% ~ 40%, then target the next 30% ~ 40% (Bailey et al., 2000; Li et al., 2011); *method 3* - preleach 60% ~ 70%, then target the next ~ 20% (Liu et al., 2013; Li et al., 2020). We conducted a four-step leaching experiment (simplified from the ten-step leaching method above), which was designed to dissolve ~30% carbonate in each step after a 1 M ammonium acetate prewash. Accordingly, the leachates from steps 1, 2 and 3 roughly correspond to cut-offs in methods 1, 2, and 3 respectively, which were then measured for Sr isotopes.

3.2. Elemental analyses

Elemental analyses, including major, trace elements and REY analysis, were carried out in the Cross-Faculty Elemental Analysis Facility, University College London (UCL), for solutions of bulk carbonate and all leaching supernatants of selected samples and rock standards. For bulk carbonate measurement, sample powder (~10-20 mg) was dissolved in 5ml 2% HNO_3 for 24 hr, then centrifuged at 3600 rpm, for 5mins. The supernatant was collected and diluted with 2% HNO_3 for elemental analysis. Major elements including Ca, Mg, K, Na, Fe, Mn, Al, and Sr were measured by inductively coupled plasma optical emission spectrometry (Varian 720 ICP-AES). The analyses achieved an error of < 3 % for the analyzed elements based

on long-term reproducibility of the laboratory measurement. Trace element Rb, Y and rare earth element (REE) concentrations of carbonate leachates were analysed by inductively coupled plasma mass spectrometry (Agilent 7900 ICP-MS) and run against multi-element matrix-matched standards within an appropriate concentration range. Replicate extractions gave a relative standard deviation (RSD) of less than 5%. The stepwise REY concentration is calculated based on the mass of Ca, Mg carbonate being leached out (i.e., mass of REY divided by mass of $CaCO_3$+$MgCO_3$ in each step). All REY concentrations were normalized to post-Archean Australian Shale (PAAS: Pourmand et al., 2012).

3.3. Sr, C, O isotope analysis

For Sr isotopes, small polypropylene columns with polypropylene frits (~30μm) and ~1cm thickness of Eichrom® Sr specific resin were used for Sr separation in an ISO 7 (Class 10000) geochemistry laboratory at UCL. The supernatant in each leaching step was dried in Teflon beakers on a hotplate before being dissolved in 0.5ml 4M nitric acid and then passed through the precleaned and conditioned columns. The collected eluant was then dried and redissolved in 2% HNO_3 for Sr isotope measurement. The $^{87}Sr/^{86}Sr$ ratios were measured using a Nu Instruments Plasma 3 multi-collector inductively coupled plasma mass spectrometer (MC-ICP-MS) at UCL. Each sample was bracketed by two measurements of NBS 987. Systematic offsets in analytical sessions were corrected by normalizing the average of the bracketing standards to a reference value (NBS 987) of $^{87}Sr/^{86}Sr = 0.710252$. The internal standard error (2se) for each sample and standard deviation (2sd) for repeated measurements of NBS 987 in each session are both reported in this study, see **Table 3 and 4**. An internal carbonate standard (N1, modern shell) was processed along with sample leachates and its multi-run average is 0.70918 (2sd =1.8×10^{-5}, n=10). A

procedural blank was included in each batch of samples, and Sr quantities (below 0.1 ng) were negligible compared to the analyte signal.

For C, O isotope analysis, powdered carbonate was analyzed at the Bloomsbury Environmental Isotope Facility (BEIF) at University College London on a continuous-flow (ThermoFisher Delta V) mass spectrometer linked to a Gas Bench II device. External error (1σ) from standards (NBS19) was better than ± 0.04 ‰ for $\delta^{13}C$ and ± 0.05‰ for $\delta^{18}O$. All values are reported using the Vienna Pee Dee Belemnite notation (VPDB) relative to NBS19 ($\delta^{13}C$=1.95‰, $\delta^{18}O$=-2.2‰).

4. Results

4.1. Step leaching Sr isotopic and elemental variation

$^{87}Sr/^{86}Sr$ and elemental (Mg/Ca, Mn/Sr, Sr/Ca, Rb/Sr, Ba/Ca, K/Ca, Al/Ca, Fe/Ca) ratios, and cumulative % carbonate dissolution for all leaching steps for samples (J95, G99, K46, G90, HXSZ4) and two rock standards (LS19, 88a) are shown in **Table 3, Fig. 1-4.**

For all types of rocks, $^{87}Sr/^{86}Sr$ ratios in step 0 (1 M NH_4AC prewash) show very high (or the highest) values. The pure limestone sample (LS19) exhibits lowest Sr isotope ratios through the three middle leachates (S1-S3), corresponding to the middle ~5%-70% carbonate dissolution, before rising in the final two leaching steps. The lowest Sr isotope values in S1-S3 are associated with lower Mg/Ca, Mn/Sr, Rb/Sr, K/Ca, Ba/Ca, Al/Ca, Fe/Ca ratios, and a higher Sr/Ca ratio than both the prewash and the last two leachates. The middle three values of LS19 measured in this study (0.70543 ± 0.00003) are consistent with what has been reported (0.705439 ± 0.000003) by Zhou et al., (2020) within error.

Sequential dissolution of argillaceous and dolomitic limestone (SDL and HDL), following the NH_4Ac prewash, results in a minimum $^{87}Sr/^{86}Sr$ value in S1 (c. 10% ~ 30% carbonate dissolution), an increase in S2 and a dramatic increase in the last two steps. The lowermost Sr isotope ratios of argillaceous and dolomitic samples range from 0.70503 to 0.70533, similar to the reported seawater range during 1.60-1.55 Ga (0.7046-0.7056; e.g., Ray et al., 2003; Kuznetsov et al., 2008; Bellefroid et al., 2018; Tan et al., 2020). In general, Sr isotope ratios are seen to be higher when elemental ratio indicators of alteration, such as Mg/Ca, Rb/Sr, Ba/Ca, and K/Ca, are higher.

From S1 onwards, Sr/Ca shows a gradually decreasing trend, while Mn/Sr shows an increasing trend. Al/Ca and Fe/Ca ratios rise more clearly only after S6.

The $^{87}Sr/^{86}Sr$ values of the pure dolostone sample (88a) have two apparent dips at S1-S2 and S7-S8, which correspond to 20% ~ 30% and 70% ~ 80% of carbonate dissolution, respectively, with the lowest value found in S7-S8 (0.710043) **(Fig. 4A)**. Mg/Ca increases gradually from S0 to S9. Mn/Sr, Sr/Ca, Rb/Sr, Ba/Ca and K/Ca ratios decrease after S0 and remain stable since S1, which are mirrored by Al/Ca, Fe/Ca ratios.

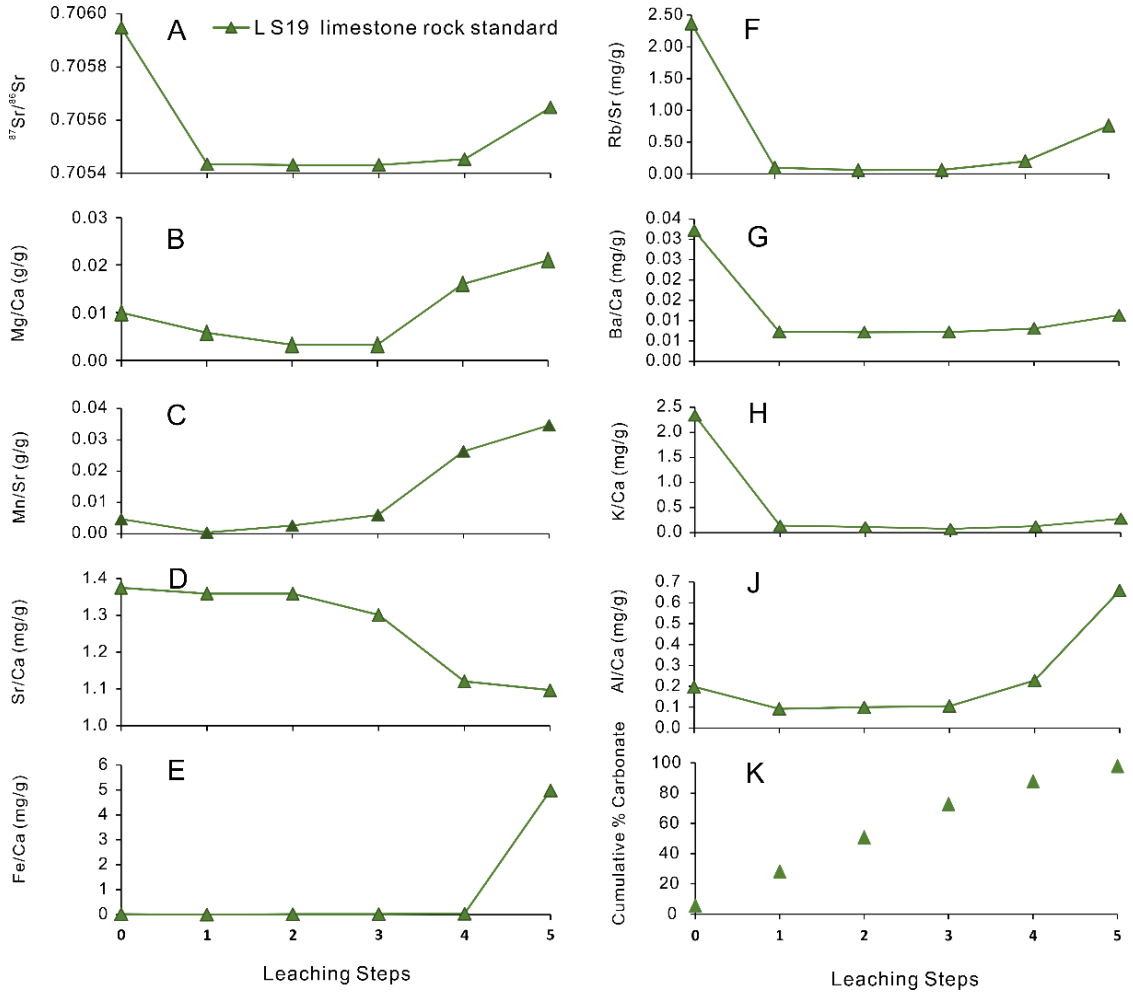

Fig. 1. Leaching pattern of pure limestones (LST) standard LS19

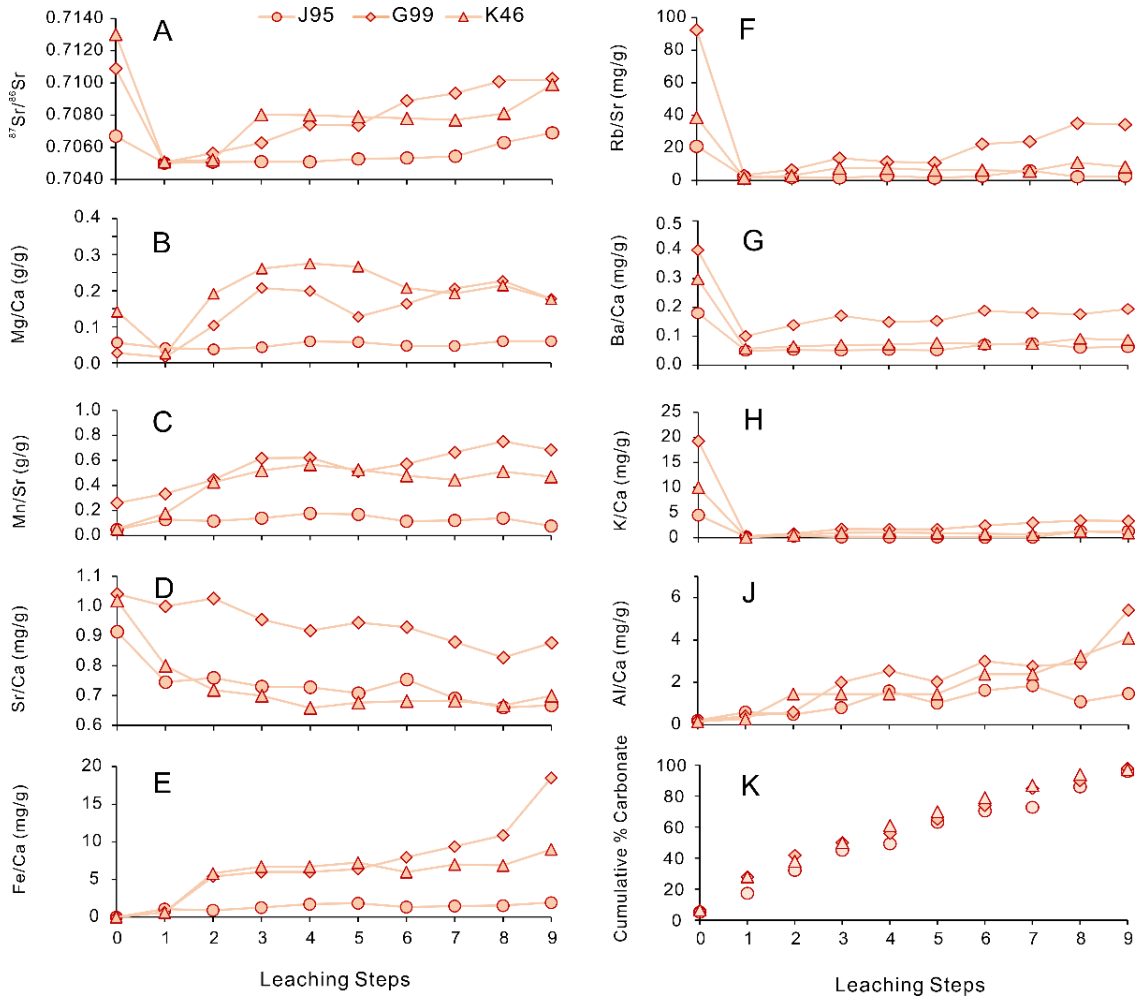

Fig. 2. Leaching pattern of argillaceous and slightly dolomitic limestones (SDL)

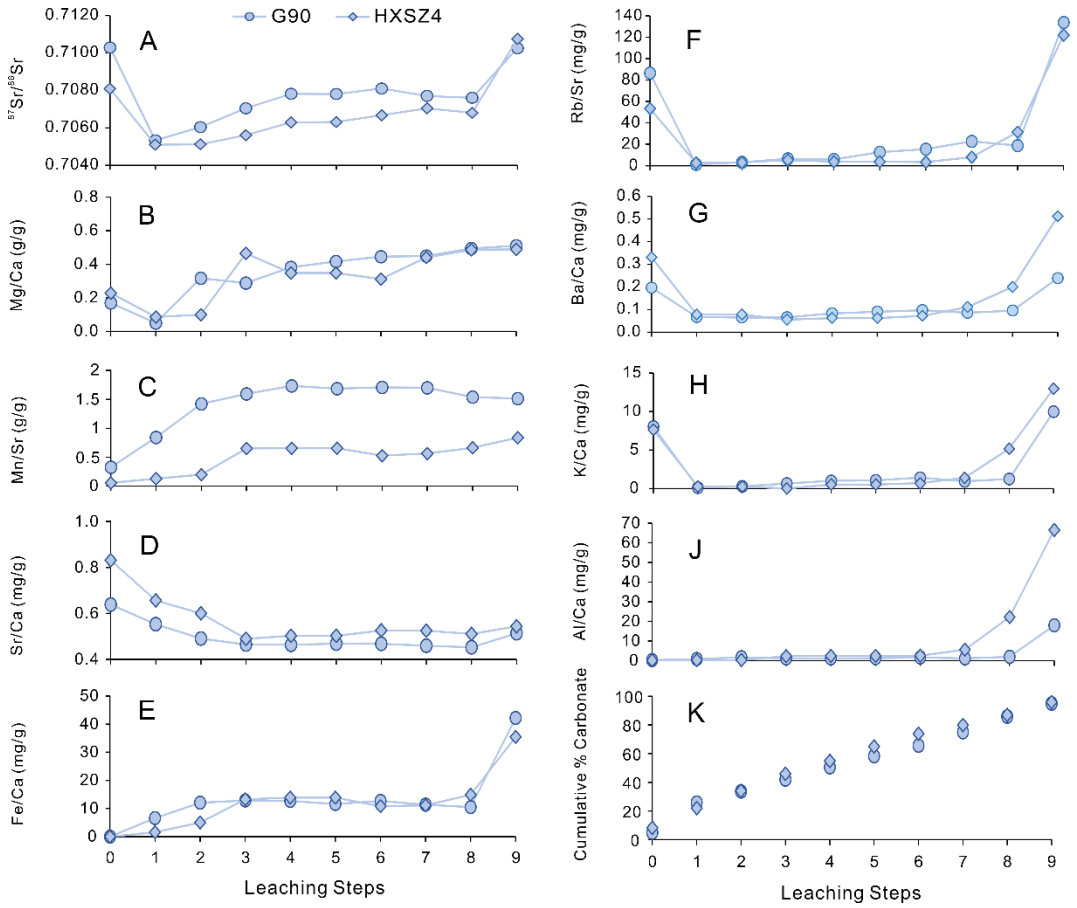

Fig. 3. Leaching pattern of argillaceous and highly dolomitic limestones (HDL).

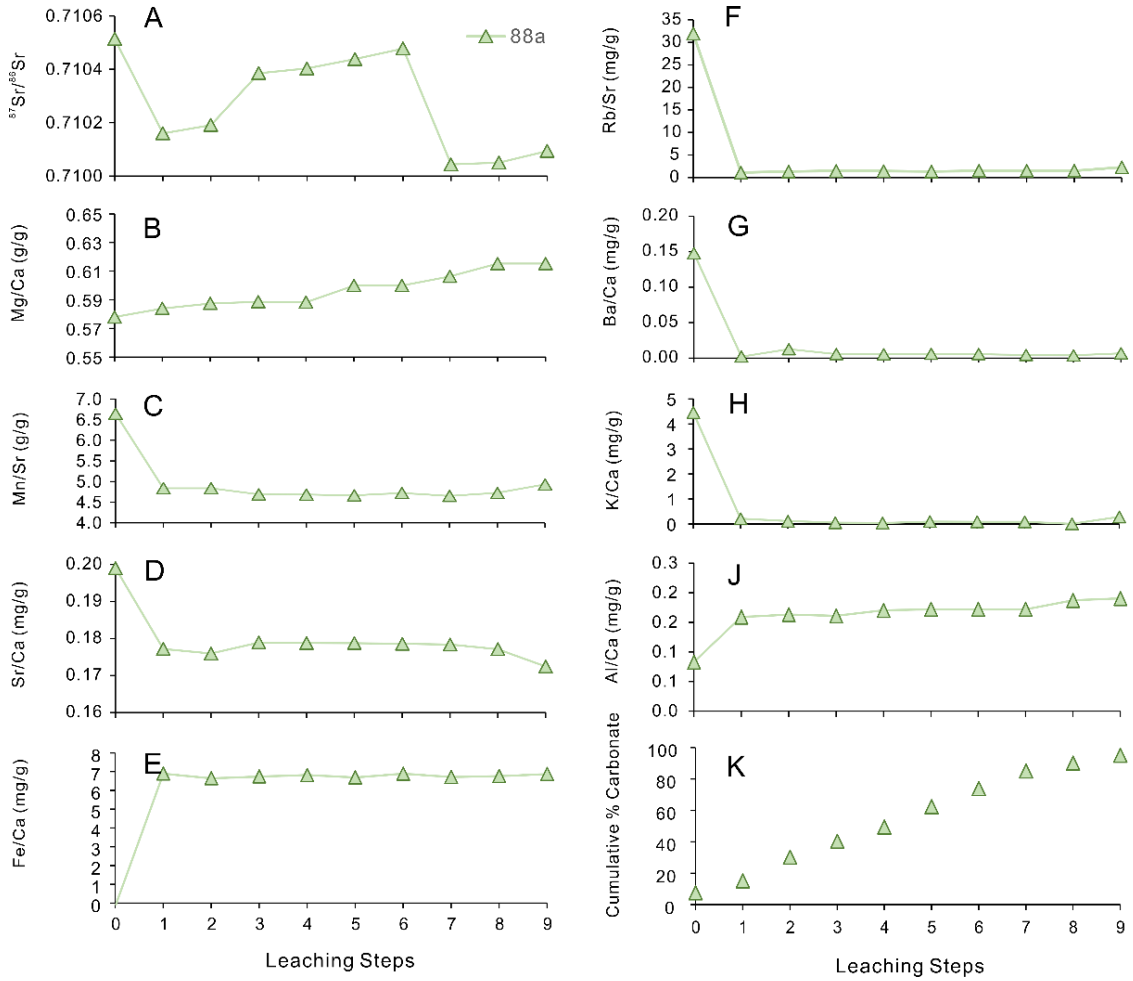

Fig. 4. Leaching pattern of pure dolostone rock standard 88a (DST).

289

Sample	Carbonate dissolved (cumulative)	$^{87}Sr/^{86}Sr$	2se	2sd	Sample	Carbonate dissolved (cumulative)	$^{87}Sr/^{86}Sr$	2se	2sd
LS19-0	5.64%	0.70595	2.54E-05	3.07E-05	G90-0	5.02%	0.71029	2.32E-05	3.03E-05
LS19-1	28.34%	***0.70543***	2.00E-05	3.07E-05	G90-1	25.93%	***0.70533***	2.18E-05	3.03E-05
LS19-2	50.84%	***0.70543***	2.08E-05	3.07E-05	G90-2	33.77%	0.70580	2.18E-05	3.03E-05
LS19-3	72.89%	***0.70543***	2.30E-05	3.07E-05	G90-3	41.96%	0.70704	1.95E-05	3.03E-05
LS19-4	87.93%	0.70545	2.20E-05	3.07E-05	G90-4	50.69%	0.70781	2.50E-05	3.03E-05
LS19-5	98.00%	0.70565	2.54E-05	3.07E-05	G90-5	58.38%	0.70780	2.46E-05	3.03E-05
J95-0	5.16%	0.70669	2.72E-05	3.07E-05	G90-6	65.64%	0.70810	1.92E-05	3.03E-05
J95-1	17.34%	***0.70503***	2.10E-05	3.07E-05	G90-7	75.14%	0.70770	2.32E-05	3.03E-05
J95-2	32.20%	0.70509	2.32E-05	3.07E-05	G90-8	86.00%	0.70760	1.84E-05	3.03E-05
J95-3	40.10%	0.70511	2.04E-05	3.07E-05	G90-9	95.00%	0.71025	2.38E-05	3.03E-05
J95-4	49.29%	0.70510	1.77E-05	3.07E-05	HXSZ4-0	8.35%	0.70808	2.38E-05	3.03E-05
J95-5	63.12%	0.70529	2.08E-05	3.07E-05	HXSZ4-1	22.00%	***0.70520***	2.56E-05	3.03E-05
J95-6	70.49%	0.70534	1.68E-05	3.07E-05	HXSZ4-2	34.00%	***0.70522***	2.40E-04	3.03E-05
J95-7	77.78%	0.70544	2.14E-05	3.07E-05	HXSZ4-3	46.00%	0.70561	2.00E-05	3.03E-05
J95-8	86.00%	0.70630	2.08E-05	3.07E-05	HXSZ4-4	55.00%	0.70629	2.06E-05	3.03E-05
J95-9	96.00%	0.70690	1.93E-05	3.07E-05	HXSZ4-5	65.00%	0.70631	1.80E-05	3.03E-05
G99-0	6.09%	0.71089	2.26E-05	3.07E-05	HXSZ4-6	74.00%	0.70667	1.90E-05	3.03E-05
G99-1	28.00%	***0.70508***	2.52E-05	3.07E-05	HXSZ4-7	80.00%	0.70704	1.82E-05	3.03E-05
G99-2	41.88%	0.70525	2.96E-05	3.07E-05	HXSZ4-8	87.00%	0.70680	2.50E-05	3.03E-05
G99-3	50.00%	0.70628	2.72E-05	3.07E-05	HXSZ4-9	96.00%	0.71075	2.82E-05	3.03E-05
G99-4	56.00%	0.70741	2.48E-05	3.07E-05	88-0	1.10%	0.71051	2.46E-05	3.03E-05
G99-5	65.00%	0.70737	2.04E-05	3.07E-05	88-1	13.87%	0.71016	1.38E-05	3.03E-05

G99-6	74.00%	0.70889	2.00E-05	3.07E-05	88-2	29.05%	0.71016	2.12E-05	3.03E-05
G99-7	85.00%	0.70936	2.18E-05	3.07E-05	88-3	39.64%	0.71039	2.56E-05	3.03E-05
G99-8	90.00%	0.71017	2.32E-05	3.07E-05	88-4	48.63%	0.71040	2.50E-05	3.03E-05
G99-9	98.00%	0.71017	2.90E-05	3.07E-05	88-5	61.29%	0.71044	2.56E-05	3.03E-05
K46-0	6.51%	0.71301	2.00E-05	3.07E-05	88-6	73.36%	0.71048	1.81E-05	3.03E-05
K46-1	28.00%	***0.70507***	2.96E-05	3.07E-05	88-7	83.97%	***0.71004***	2.38E-05	3.03E-05
K46-2	38.00%	0.70522	2.20E-05	3.07E-05	88-8	90.63%	***0.71005***	2.78E-05	3.03E-05
K46-3	50.00%	0.70803	2.36E-05	3.07E-05	88-9	93.81%	0.71009	2.30E-05	3.03E-05
K46-4	61.00%	0.70800	2.92E-05	3.07E-05					
K46-5	70.00%	0.70790	2.42E-05	3.07E-05					
K46-6	79.00%	0.70780	2.74E-05	3.07E-05					
K46-7	87.00%	0.70770	2.10E-05	3.07E-05					
K46-8	94.00%	0.70810	2.12E-05	3.07E-05					
K46-9	97.00%	0.70990	2.08E-05	3.07E-05					

Table 3. $^{87}Sr/^{86}Sr$ and cumulative carbonate dissolution in each step of each sample. The internal standard error (2se) for each sample and standard deviation (2sd) for repeat measurements of NBS 987 in each session are reported here. The lowest value (within 2sd) for each sample is highlighted by ***bold and italic***. See **Table S1 in Supplementary Materials** for major and trace element data.

4.2. Step leaching REY pattern

To help understand the leaching pattern of Sr isotopes for argillaceous and dolomitic limestones, we also examined their step-leaching REY pattern **(Fig. 5, Table S2 in Supplementary Materials).** Supernatants of ammonium acetate prewash exhibit a non-seawater pattern, either being flat or displaying a positive Eu anomaly. Ba/Sm ratios show a strong correlation with Eu_{sn}/Eu^*_{sn} for all argillaceous and dolomitic limestones, strongly implying that Eu anomalies in the S0 resulted from BaO interference in the ICP-MS (Jarvis et al., 1989), while no such correlation was observed for other leaching steps **(Fig. 6**). A seawater-like REY pattern occurs in S1 for almost all samples but tends to be flat in the subsequent leaching steps. The only exception is G99, which shows a non-seawater pattern and a negative Eu anomaly throughout all leaching steps. The Y/Ho ratios of the argillaceous and dolomitic limestone in each leaching step are shown in **Fig. 7.** All samples display a Y/Ho ratio >36 in S1, and then the ratio gradually decreases in subsequent steps.

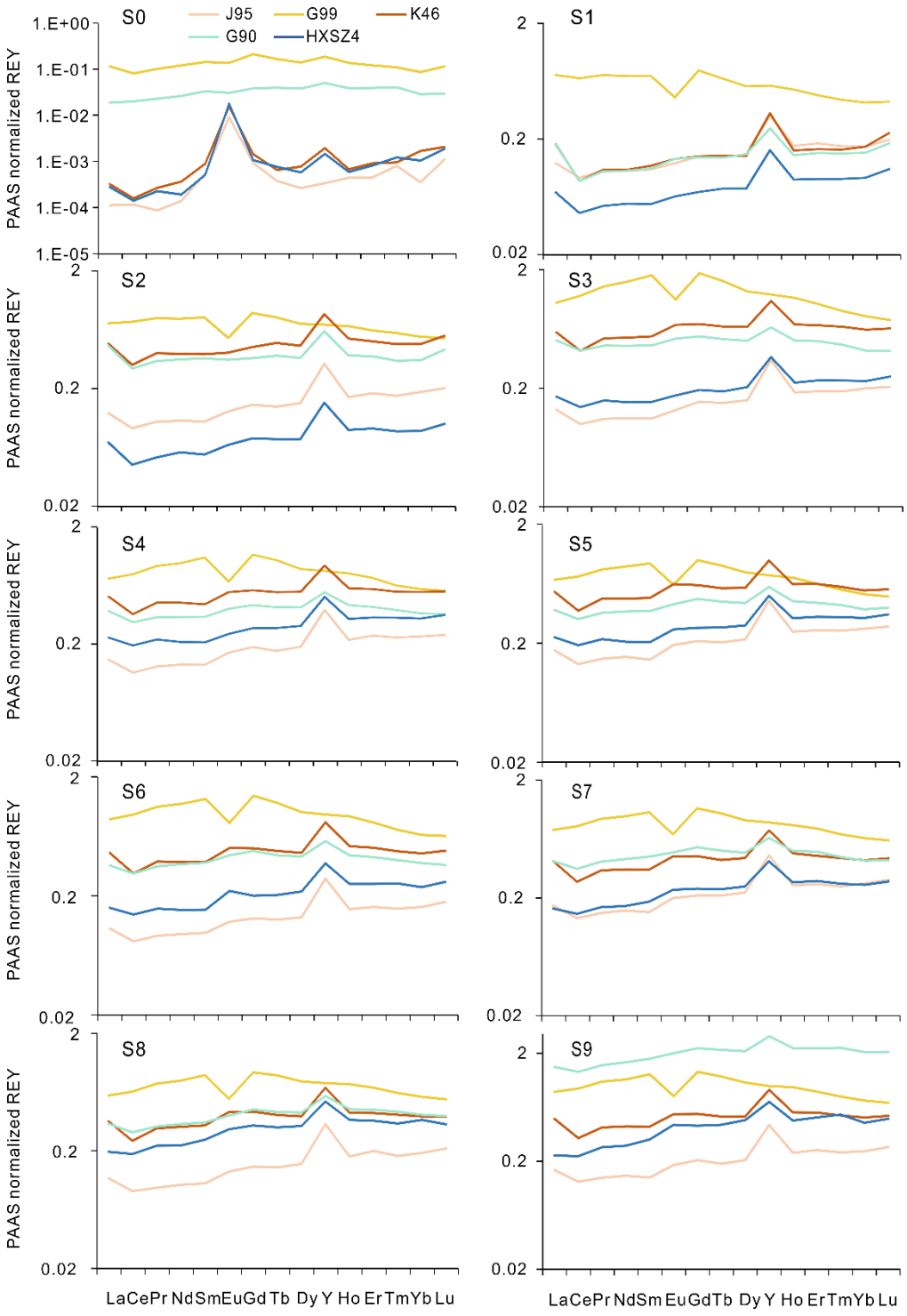

Fig. 5. Ten-step REY leaching pattern for argillaceous and dolomitic limestones. Each plot shows each step, and lines with different colours represent different samples. All plots are in log scale. Except for G99, all samples show the seawater pattern in S1 (some also in S2).

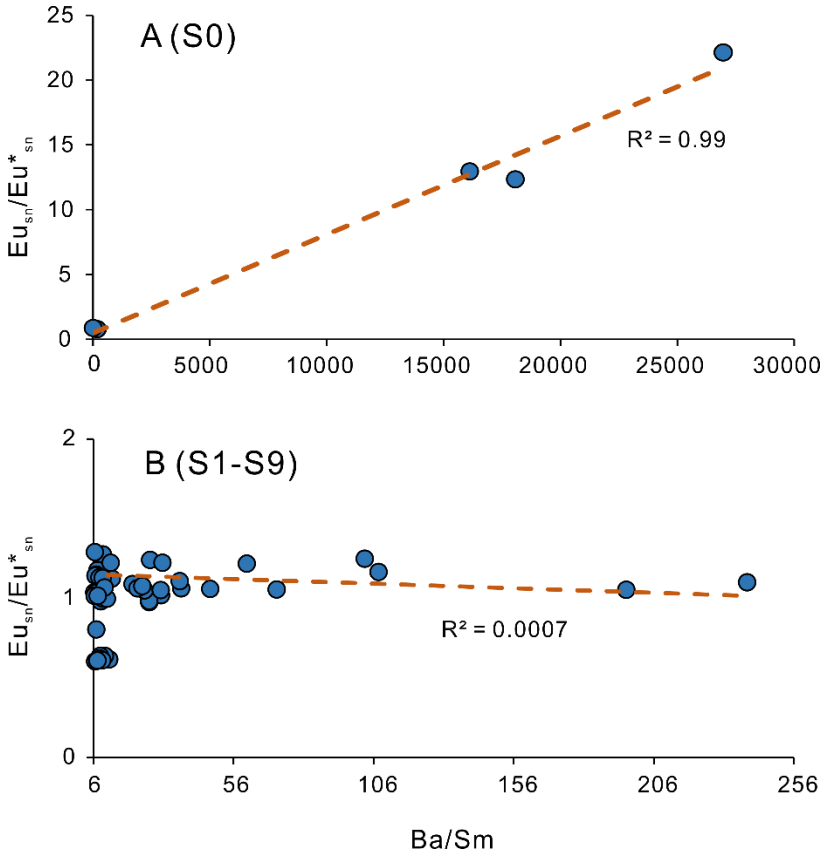

Fig. 6. Crossplots of Ba/Sm and Eu_{sn}/Eu^*_{sn} for 5 argillaceous and dolomitic limestones. A). Crossplot of Ba/Sm and Eu_{sn}/Eu^*_{sn} for S0 (NH_4Ac prewash) shows a strong correlation; B). Crossplot of Ba/Sm and Eu_{sn}/Eu^*_{sn} for S1-S9 shows no correlation.

$Eu_{sn}/Eu^*_{sn} = 2*[Eu]_{sn}/([Sm]_{sn}+[Gd]_{sn})$; sn: shale normalized.

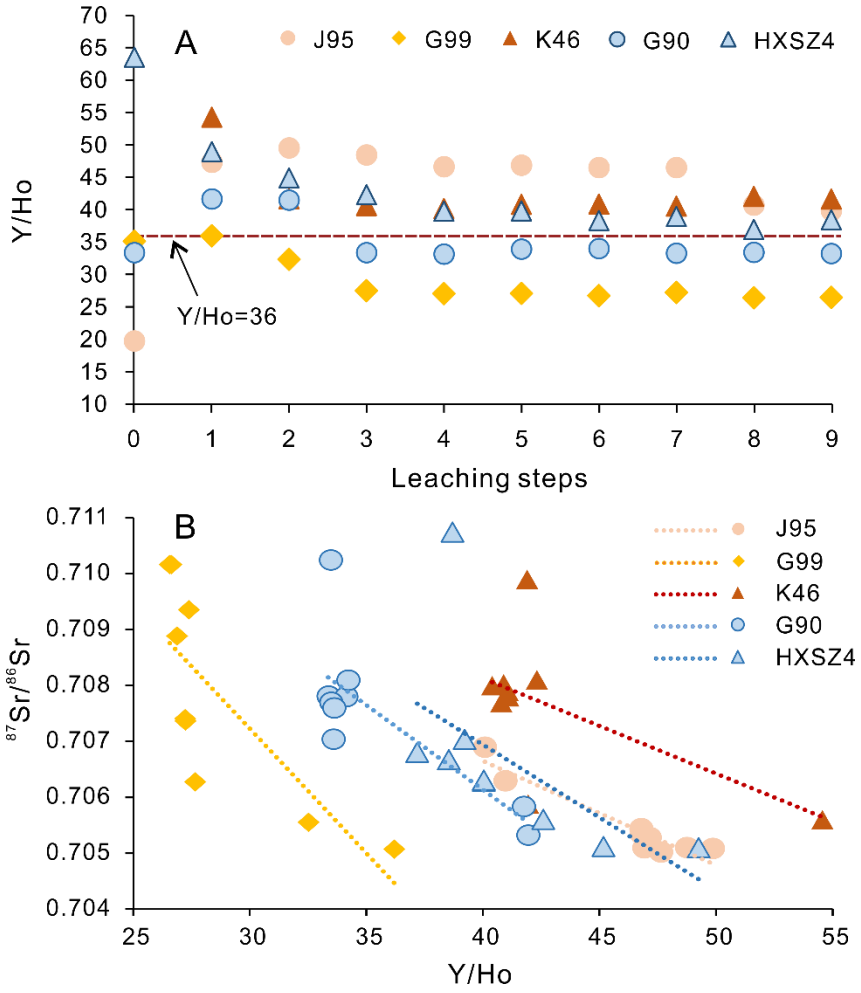

Fig.7. A). Y/Ho ratios in each leaching step of each sample. All samples in S1 show Y/Ho > 36 (seawater signal; less clay contamination). B). Cross-plot of Sr isotopes versus Y/Ho in each leaching step of each sample. A negative correlation was observed for each sample.

4.3. Leaching methods' comparison results

As mentioned before (section 3.1.2), 20 carbonate samples with various purity are selected to test the three different leaching methods: *Method 1* - target the first 10% ~ 30% after NH_4Ac prewash (this study); *Method 2* - preleach 30%, target the next 30% (Bailey et al., 2000; Li et al., 2011); *Method 3* - preleach 60%-70%, target the next ~20% (Liu et al., 2013; Li et al., 2020). Sr isotope results and bulk carbonate information for each sample are shown in **Table 4** and **Fig. 8.** For pure limestone

311 (HXS1), no significant differences are evident between the three different leaching
312 cut-offs (i.e., the disparities are within measurement errors). For some very
313 argillaceous samples (carbonate content ⩽ ~50%) such as G94 and G97, applying
314 higher leaching cut-offs (i.e., methods 2 and 3) would increase the Sr isotope ratios
315 dramatically (up to 0.009 compared with targeting the first 10% ~ 30% after
316 prewash). Relatively pure dolomitic limestones (% carbonate > 80% and 0.025 <
317 Mg/Ca < 0.6) also tend to reach the lowest values using method 1, the method this
318 study recommends, although adjusting cut off points would not produce the more
319 extreme differences seen in very argillaceous samples. When Mg/Ca > 0.4, the
320 situation becomes quite complicated, i.e., the lowest value could be generated by any
321 of the leaching methods, but no near-seawater values are observed.

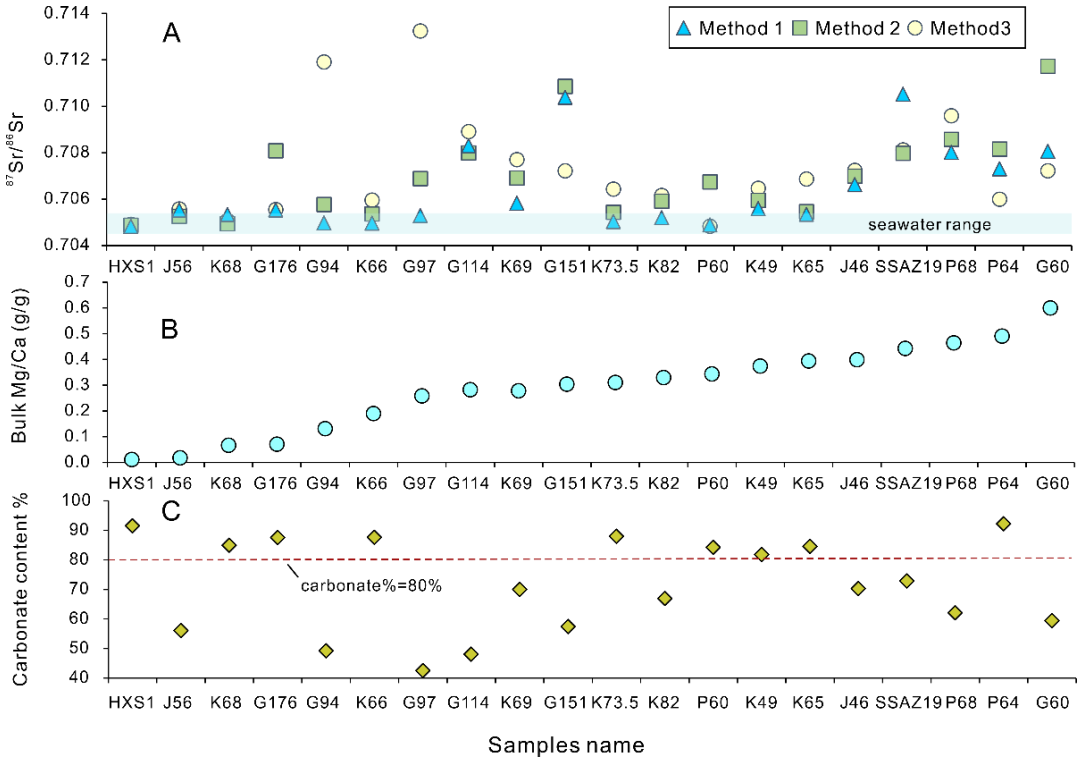

Fig. 8. A). Sr isotope results by applying different leaching Methods to each sample. Blue horizontal bar presents the possible seawater range from ~0.7046 to ~0.7056 of the Member III of Gaoyuzhuang Fm (see section 2.1 for details). **B).** Bulk Mg/Ca ratios of samples range

from 0.01 to 0.6. **C).** Carbonate content of each sample. Redline shows % carbonate = 80%, which divides samples into high purity (> 80%) and low purity (< 80%).

Samples	$^{87}Sr/^{86}Sr$			Bulk rock			
	Method 1	*Method 2*	*Method 3*	Carbonate content (%)	Mg/Ca (g/g)	Mn/Sr (g/g)	[Sr/(Ca+Mg)] (ppm)
HXS1	***0.70487 (2)***	***0.70488 (2)***	***0.70488 (2)***	91.49	0.01	0.00	8295.27
J56	***0.70523 (2)***	***0.70526 (3)***	0.70558 (2)	56.06	0.02	0.31	579.20
K68	***0.70493 (2)***	***0.70495 (1)***	0.70508 (2)	84.95	0.07	0.35	521.65
G176	***0.70653 (2)***	0.70809 (2)	***0.70655 (2)***	87.55	0.07	3.04	192.74
G94	***0.70498 (1)***	0.70577 (3)	0.71191 (2)	49.20	0.13	0.23	1106.69
K66	***0.70496 (2)***	0.70536 (2)	0.70596 (2)	87.62	0.19	0.30	662.06
G97	***0.70509 (2)***	0.70689 (2)	0.71325 (3)	49.52	0.26	0.48	662.05
G114	0.70829 (2)	***0.70800 (2)***	0.70891 (2)	48.06	0.28	2.22	404.09
K69	***0.70583 (3)***	0.70692 (2)	0.70771 (2)	70.00	0.28	0.40	779.79
G151	*0.71038 (3)*	0.71086 (3)	***0.70723 (2)***	57.45	0.30	3.82	233.93
K73.5	***0.70503 (2)***	***0.70505 (1)***	0.70643 (2)	88.00	0.31	0.44	842.47
K82	***0.70520 (2)***	0.70592 (2)	0.70616 (1)	66.94	0.33	0.56	442.54
P60	***0.70483 (2)***	0.70674 (1)	***0.70484 (2)***	84.20	0.34	1.69	218.06
K49	***0.70559 (2)***	0.70595 (2)	0.70647 (2)	81.80	0.37	0.84	389.22
K65	***0.70535 (1)***	0.70546 (2)	0.70687 (2)	84.59	0.39	0.67	499.80
J46	***0.70662 (2)***	0.70700 (2)	0.70724 (2)	70.28	0.40	4.43	261.99
SSAZ19	0.71053 (2)	***0.70798 (1)***	0.70812 (1)	72.86	0.44	2.00	301.35
P68	***0.70802 (2)***	0.70857 (2)	0.70959 (2)	62.13	0.46	3.34	178.26
P64	0.70731 (2)	0.70816 (1)	***0.70600 (3)***	92.21	0.49	2.89	145.95
G60	0.70806 (2)	0.71173 (2)	***0.70723 (1)***	59.46	0.60	4.29	124.37

Table 4. $^{87}Sr/^{86}Sr$ results of applying different cut-offs to samples with various bulk rock information. Note that the number in the bracket represents the 2se in the least digit. ***Bold and italic*** highlight the lowest value (within 2sd) among three different leaching methods (2sd = 3.11E-05).

5. Discussion

5.1. Explanations of step-leaching Sr isotope and elemental variations

The leaching step showing the lowest, and likely most pristine Sr isotope ratios may correspond to the effective isolation of either the clay-free carbonate fraction and/or the least altered carbonate phase in a given sample. Rb, K, and Al concentrations are typical indicators of aluminosilicate Sr contamination (Banner et al., 1988; McArthur, 1994; Montañez et al., 1996). Rb and K can both be used to track clay surface-bound Sr released by ion exchange as well as detrital / authigenic clay dissolution. By contrast, Al might not be a suitable proxy for Sr released from clay surfaces due to its insoluble nature (Bellefroid et al., 2018), but it is a strong sign of aluminosilicate dissolution (Wierzbowski et al., 2012). Mg/Ca is used to quantify the relative contribution of calcite and dolomite during leaching steps and the degree of dolomitization in a sample. Mn/Sr, Fe/Ca and Sr/Ca are often used as indices of alteration, and Mn/Sr, Fe/Ca are generally expected to be higher, while Sr/Ca is considered to be lower in diagenetically altered samples than in coeval seawater (Banner and Hanson, 1990; Gorokhov et al., 1995; Kaufman and Knoll, 1995). This general relationship is complicated by variable redox conditions, diagenetic fluids, and mineralogy. For instance, compared with calcite, dolomite generally has a greater preference for Fe and Mn (Mazzullo, 1992) and a lower preference for Sr (Vahrenkamp and Swart, 1990). Apart from diagenetic phases, the stepwise leaching patterns of Mn/Sr and Fe/Ca might also indicate the dissolution of non-carbonate phases such as Fe-Mn oxides (e.g., Zhang et al., 2015)

5.1.1. Argillaceous and dolomitic limestones

5.1.1.1. Similarities among samples

All argillaceous and dolomitic limestones exhibit a similar Sr isotope leaching pattern, i.e., reaching a nadir in S1, and then rising through subsequent steps. A comparable pattern was previously reported by Bellefroid et al., (2018) on limestones of the Dhaiqa and Tieling formations, where it was referred to as a "V"-shaped pattern. The extremely radiogenic Sr and high Rb/Sr, K/Ca ratios of the first leaching step (S0) by ammonium acetate are contributed to significantly by Sr in ion exchange sites in clay minerals and trace metals adsorbed on mineral surfaces (Morton, 1985; Gao, 1990; Bailey et al., 2000). The dramatic drop of Rb/Sr, K/Ca and $^{87}Sr/^{86}Sr$ in the following step (S1, in some cases S2) indicates that pre-cleaning has effectively removed this weakly surface-bound Sr. The increase in Rb/Sr, K/Ca, and Al/Ca ratios and more radiogenic $^{87}Sr/^{86}Sr$ from S2, and especially after S6, most likely implies partial dissolution of residual aluminosilicate, considering the samples' argillaceous lithology. Intriguingly, we found that Ba/Ca follows a similar pattern and exhibits a strong linear correlation with Rb/Sr (R^2 >0.95; **Fig. 9**). Previous research found that clay is one of the main Ba-carriers in marine sediments (Rutten and de Lange, 2002; Gonneea and Paytan, 2006). Therefore, the strong correlation between Rb and Ba might support the use of Ba/Ca ratios as indicators of clay contamination. Low Mg/Ca ratios (all below 0.1) in S1 for all argillaceous and dolomitic limestones indicate that the calcite proportion was leached out before dolomite as calcite reacts much faster with acid than dolomite. Gradually increasing Mg/Ca, Mn/Sr and Fe/Ca and decreasing Sr/Ca after S1 demonstrate that the calcite proportion in S1 is the closest to primary carbonate phase, which is released before secondary calcite, dolomite and non-carbonate minerals are dissolved in subsequent leaching steps.

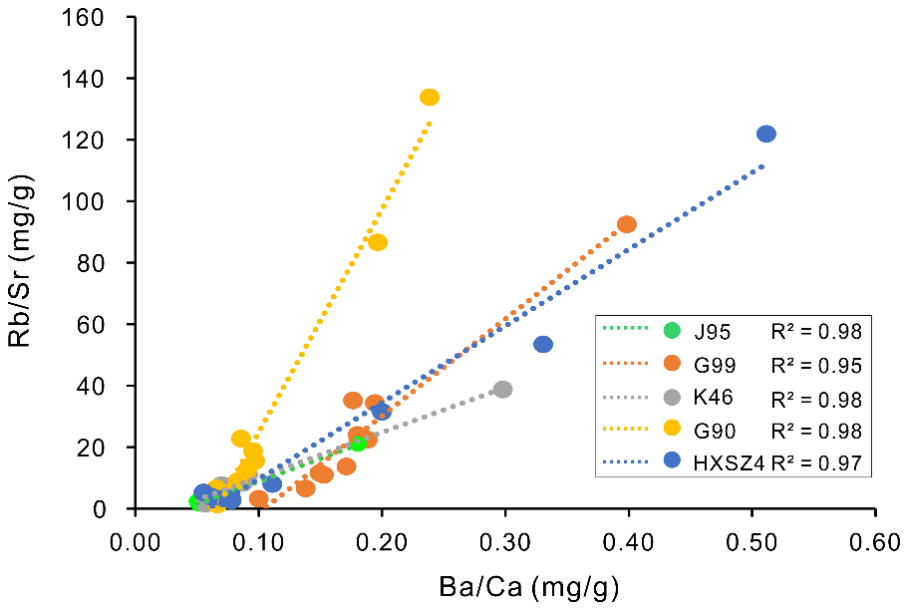

Fig. 9. Cross-plot of stepwise Rb/Sr versus Ba/Ca ratios for each argillaceous and dolomitic limestone. The correlation coefficient (R^2) of each sample is > 0.95.

5.1.1.2. Disparities among samples

Slight differences still exist among samples, and more explanations could be investigated by cross plotting the stepwise elemental ratios (Rb/Sr, Mn/Sr and Mg/Ca) against Sr isotopes **(Fig. 10)**. After reaching the minimal value in S1, J95, a slightly dolomitic limestone, shows a slower rebound to a higher value compared with other samples. The step leach Rb/Sr, Mn/Sr and Mg/Ca ratios of J95 show the weakest relationship with Sr isotopes (from S2 to S9) compared with other samples, which indicates that clay contamination and leaching of diagenetic phases did not influence the Sr isotope values immediately from S2 **(Fig. 10. A1-A3)**. With a Mg/Ca ratio (0.06) close to limestone (< 0.025), and a carbonate content of ~70%, J95 contains a relatively higher proportion of "clean and primary" calcite compared with the other four argillaceous and dolomitic limestones (G99, K46, G90, HXSZ4). In contrast, all samples except J95 show positive correlations between Rb/Sr and $^{87}Sr/^{86}Sr$ (**Fig. 10. B1-E1**), whereby the 'dirtiest' sample (G99) has the strongest correlation (R^2=0.91, **Fig. 10. B1**). This result demonstrates that for very argillaceous samples, the Sr

isotope leaching pattern is strongly and rapidly influenced by the dissolution of clay minerals, immediately following the first leachate, even when using a weak acid. Therefore, caution needs to be taken to avoid over-leaching when dealing with argillaceous samples, while the first 10% ~ 30% after prewash would seem to represent the cleanest portion. One possible factor that influences clay mineral dissolution could be reaction time. The longer agitation in ultra-sonic bath and reaction time since S3 **(Table 2)** could have contributed to the dissolution of more clay minerals. This may indicate that when leaching dirty/muddy samples for Sr isotopes analysis, shorter reaction time is preferable. Cross plots of stepwise Mg/Ca and Mn/Sr versus Sr isotope ratios (**Fig. 10. A2-E2, A3-E3**) in most cases exhibit a stronger covariation in highly dolomitic limestone (G90, HXSZ4) than for slightly dolomitic limestones (J95, G99, K46), which possibly indicates that dissolution of the dolomitized (likely more diagenetically altered) proportion is a more important contributing factor that leads to increased $^{87}Sr/^{86}Sr$ ratios after S1 for samples with a higher degree of dolomitization.

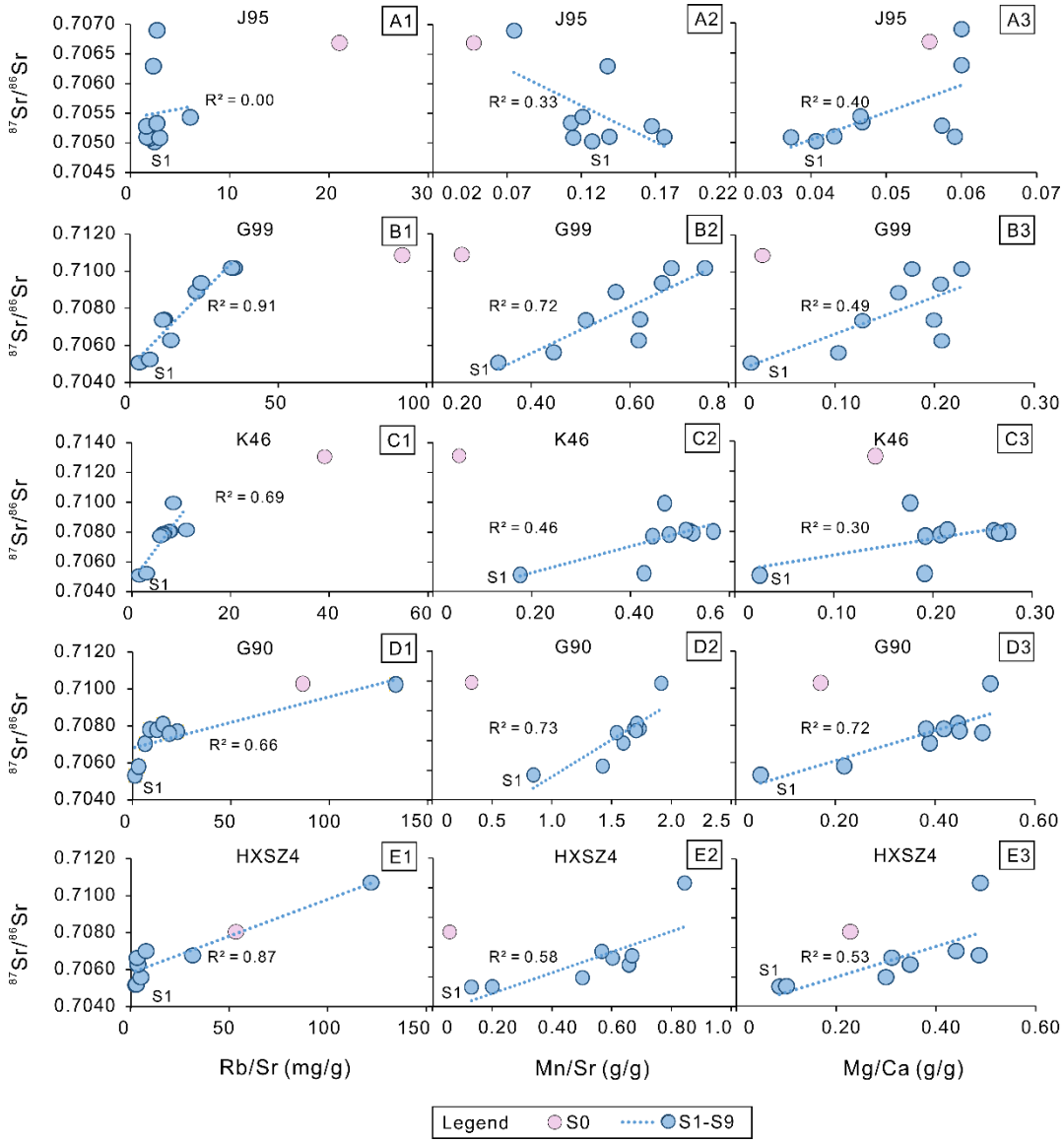

Fig. 10. Cross-plots of stepwise Rb/Sr, Mn/Sr and Mg/Ca ratios versus Sr isotopes for each argillaceous and dolomitic limestone. After S0 (prewash), correlations of Rb/Sr, Mn/Sr and Mg/Ca ratios with $^{87}Sr/^{86}Sr$ from S1 to S9 are shown in blue lines.

399 5.1.2. Two rock standards

400 Although a detailed discussion on SIS leaching methods for pure limestone and
401 dolostone samples is beyond the scope of this study, we wish to show how different
402 types of rocks behave by briefly demonstrating the leaching patterns of two rock
403 standards here. In pure limestone internal standard LS19, the “clean and primary”
404 proportion is much higher than in argillaceous and dolomitic limestones, as evidenced

by a consistent Sr isotope nadir alongside the lowest Mg/Ca and Mn/Sr ratios from S1 to S3, which corresponds to ~5% to 70% carbonate dissolution. This pattern is consistent with previous studies on limestone samples (e.g., Bailey et al., 2000; Bellefroid et al., 2018). Interestingly, the $^{87}Sr/^{86}Sr$ leaching pattern for CRM 88a exhibits two low points in steps S1/S2, and S7/S8, respectively, although none of these show seawater values as both sets of values are considerably higher than contemporaneous seawater. Therefore, the leachates of CRM 88a might contain calcite and dolomite formed at different stages of recrystallization. Nevertheless, the lowest $^{87}Sr/^{86}Sr$ value of 88a from this study is significantly lower than reported by Stammeier et al., (2020), in which the bulk sample was dissolved in 3 M HNO_3.

5.2. Comparison of $^{87}Sr/^{86}Sr$ with REY step-leaching pattern

5.2.1. Clay contamination

The high sensitivity to clay contamination of both REY patterns and Sr isotopes allows us to combine them to examine the validity of the leaching method for argillaceous and dolomitic limestones. The step leaching REY patterns of all samples, except for G99, show a seawater pattern in S1 (and some in S2), which is in line with Sr isotope leaching patterns (lowest/seawater $^{87}Sr/^{86}Sr$ values in S1). This result is also consistent with previous studies such as Tostevin et al., (2016) and Cao et al., (2020), which proposed that the early leaches contain more pristine seawater REY signals for partially dolomitized and less pure limestones. The Y/Ho ratio is one of the most effective approaches to recognizing terrigenous influences on seawater REY distribution because Ho is scavenged two times faster than Y from the surface ocean to the deep ocean (Nozaki et al., 1997). The Y/Ho ratio of seawater is consequently almost twice that of the average upper continental crust (~26–28; Taylor and McLennan, 1981; Kamber et al., 2005). Y/Ho> 36 is a commonly used threshold

value for seawater REY signals (e.g., Tostevin et al., 2016). The high Y/Ho ratios of S1 (> 36, **Fig. 7A**), before gradually decreasing in subsequent steps, further confirms our findings based on Sr isotopes, i.e. that step 1 leaches out the most primary portion of the sample. The cross plot of stepwise $^{87}Sr/^{86}Sr$ with Y/Ho ratios of each sample shows a negative covariation (either strong or weak, **Fig. 7B**), which is within expectation as dissolution of clay minerals normally would increase $^{87}Sr/^{86}Sr$ values while decreasing Y/Ho ratios. A conflicting story (i.e., a positive correlation between $^{87}Sr/^{86}Sr$ and Y/Ho) was reported by (Verdel et al., 2018), in which the authors attribute the disparity to the progressive dissolution of different combinations of sources. Nevertheless, the high consistency between REY compositions and Sr isotope leaching pattern in this study further confirmed a similar sensitivity of both systems to clay contamination.

5.2.2. Organic matter activity and dolomitization

Even though the REY pattern is consistent with the Sr isotope leaching pattern in most samples, an exception still exists in G99, the sample with a seawater like Sr isotope value but without a seawater REY pattern. One difference between G99 and other samples is that G99 contains high levels of organic compounds (total TOC of 1.3 wt %). It seems probable to us that during early diagenesis, degradation and remineralization of organic matter released adsorbed REY into porewaters so it could be incorporated into carbonate rocks, altering the original seawater REY pattern. The potential for substantial, indirect influence from organic degradation on the original carbonate REY pattern was also proposed in a recent study by Zhang and Shields, (2022). Some research shows that organic matter preferentially absorbs LREE and then releases it at depth during remineralization (Chen et al., 2015; Meyer et al., 2021), but the understanding of REY in the biological system is still at an early stage;

thus, a range of REY patterns in organic matter might be expected (Zhang and Shields, 2022). In contrast to REY patterns, the low Sr content of organic matter means that it will have little influence on pore fluid Sr isotope composition.

In addition, the highly dolomitic limestones (G90, HXSZ4) also show marine-like REY patterns in S1. This finding is consistent with the previously proposed argument that dolomitization may not significantly alter REY patterns of carbonate rocks (Banner et al., 1988; Zhang et al., 2015), although it might change Sr isotope characteristics. Therefore, while REY and Sr isotopes in carbonates have similarities (e.g., both are vulnerable to clay contamination), other factors (e.g., organic matter, dolomitization) will have different impacts on these two systems.

5.2.3. Effectiveness of ammonium acetate prewash

Ammonium acetate prewash has been widely used for leaching protocols of carbonate rocks for different proxies such as REY, Sr isotopes or Li isotopes to remove the ion-exchangeable phase (Tessier et al., 1979; Bailey et al., 2000; Kuznetsov et al., 2010; Liu et al., 2013; Pogge Von Strandmann et al., 2013; Cui et al., 2015; Bellefroid et al., 2018; Cao et al., 2020). It was suggested that using dilute acetic acid may remove Sr contamination more effectively (Bailey et al., 2000), but for the argillaceous and dolomitic limestones, it might result in over-leach as only the first 10%-30% should be extracted based on our study. In this case, using ammonium acetate instead of acetic acid for pre-cleaning to remove adsorbed Sr and Rb would be advisable as the *p*H neutral ammonium acetate would not attack as much carbonate (less than 5%).

We agree that ammonium acetate might not remove all clay surface-bound contamination (e.g., Bailey et al., 2000), but based on major and trace element, REY and Sr isotope data (as discussed in previous sections), using an ammonium acetate

prewash is still an effective option. It is also worth noting that the reaction time may play a significant role in the effectiveness of using ammonium acetate. It was suggested that a 30-minute leaching time is sufficient to achieve maximum extraction of adsorbed REE (Moldoveanu and Papangelakis, 2013; Cao et al., 2020), and this reaction time was also applied in this study.

5.3. Suitability of different leaching cut-offs for different types of samples

The pure limestone (HXS1) shows no significant difference between the three different leaching cut-offs (**Fig. 8 and Table 4**), which is consistent with the leaching pattern of pure limestone LS19 (**Fig. 1A**). However, for very argillaceous samples (% carbonate ≤ ~50%) with Mg/Ca < 0.4, such as G94, G97, applying any higher pre-leach cut-offs would produce a sizeable error compared with using the first 10%-30% after prewash, similar to what we show for the leaching pattern of G99 (**Fig. 2A**). This is because to extract the primary signal from carbonate rocks, two requirements must be met simultaneously: "relatively clay-free" and "least altered", and the only possible proportion for argillaceous and dolomitic limestones is the first 10%-30%, which would be readily missed if samples are over-leached during pre-leach. By contrast, the different leaching cut-offs for pure samples (carbonate content > 80%) with Mg/Ca < 0.4 (e.g., K66, K73.5, **Fig. 8**) do not produce as large a difference as the more argillaceous samples. In general, most samples with Mg/Ca < 0.4 yield the lowest/seawater Sr isotopic values with the first ~30% dissolution after prewash (**Fig. 8, Table 4**), which implies that for most of the limestones and slightly dolomitic limestone samples of GYZ formation, the calcite proportion dissolved in the early stage contains the primary marine signal. However, when Mg/Ca >0.4, it is unpredictable which method will yield the lowest $^{87}Sr/^{86}Sr$ value, and all lowest values are higher than the proposed seawater value (**Fig. 8**). One possible contributing

505 factor is leaching out other calcite components, in the form of dedolomitization or
506 secondary veins and cements, especially in highly dolomitic limestones and
507 dolostones (e.g., Tostevin et al., 2016).

508 By comparison, targeting the first ~10%-30% carbonate after ammonium acetate
509 prewash appears to be the most appropriate method for SIS using a wide range of
510 carbonate, especially argillaceous and dolomitic limestones with detailed thresholds
511 described in session 5.4. The widely used cut-off for bulk carbonate (with 30%
512 preleach) is more suitable for limestones or pure dolomitic limestones than for other
513 rock types. Although our study and Liu et al., (2013, 2014) show pre-leaching of
514 60%-70% can obtain the lowest $^{87}Sr/^{86}Sr$ value for highly dolomitic limestones or
515 dolostones, our data present that the lowest Sr isotopic value of CRM 88a
516 (Mg/Ca=0.6) and other samples with Mg/Ca > 0.4 are higher than contemporaneous
517 seawater, therefore, without further tests, no recommendations can be made for these
518 types of samples.

519 **5.4. Recommended cut-offs for sample screening in SIS studies**

520 Based on our leaching tests and the application of three different methods to
521 twenty samples with different sample purity and dolomitization, we noticed that
522 samples with Mg/Ca > 0.4, Mn/Sr > 2, [Sr] < 200 ppm would be less likely to retain
523 the original seawater signal (i.e., all the lowest $^{87}Sr/^{86}Sr$ ratios from these types of
524 samples yielded from three leaching methods are higher than contemporaneous
525 seawater values). Cross-plots of Sr isotopes (lowest values among three methods)
526 versus bulk Mg/Ca, Mn/Sr, % carbonate, and Sr/(Ca+Mg) **(Fig. 11)** illustrate that
527 when Mg/Ca < 0.25, [Sr] > 400 ppm, Mn/Sr < 2, the success rate (the likelihood to
528 obtain seawater value) will be high. A previous study by Li et al., (2020) suggested
529 that samples with high purities (> 75% for limestones, > 90% for dolostones) are

more suitable for SIS. Even though sample purity influences the likelihood of obtaining seawater values if the proposed seawater Sr isotopic ratio range is valid, our data (**Fig. 8A and Fig. 11D**) show that our leaching method increases that likelihood significantly, in 4 out of 6 cases with carbonate content < 70% when other thresholds are met.

Our suggested thresholds here are based on the lowermost value among three leaching methods for the 20 selected samples from the Gaoyuzhuang Formation. The thresholds proposed from this study are in agreement broadly with previously proposed thresholds by different studies (e.g., Bartley et al., 2001; Halverson et al., 2007; Bold et al., 2016; Cox et al., 2016; Gibson et al., 2019; Zhou et al., 2020). We agree that it is unlikely to have any single criterion for the robust screening of altered samples because the post-depositional history varies from basin to basin (Bartley et al., 2001; Melezhik et al., 2001; Halverson et al., 2007), but the consistency between different research might provide a valuable reference for future Precambrian SIS studies.

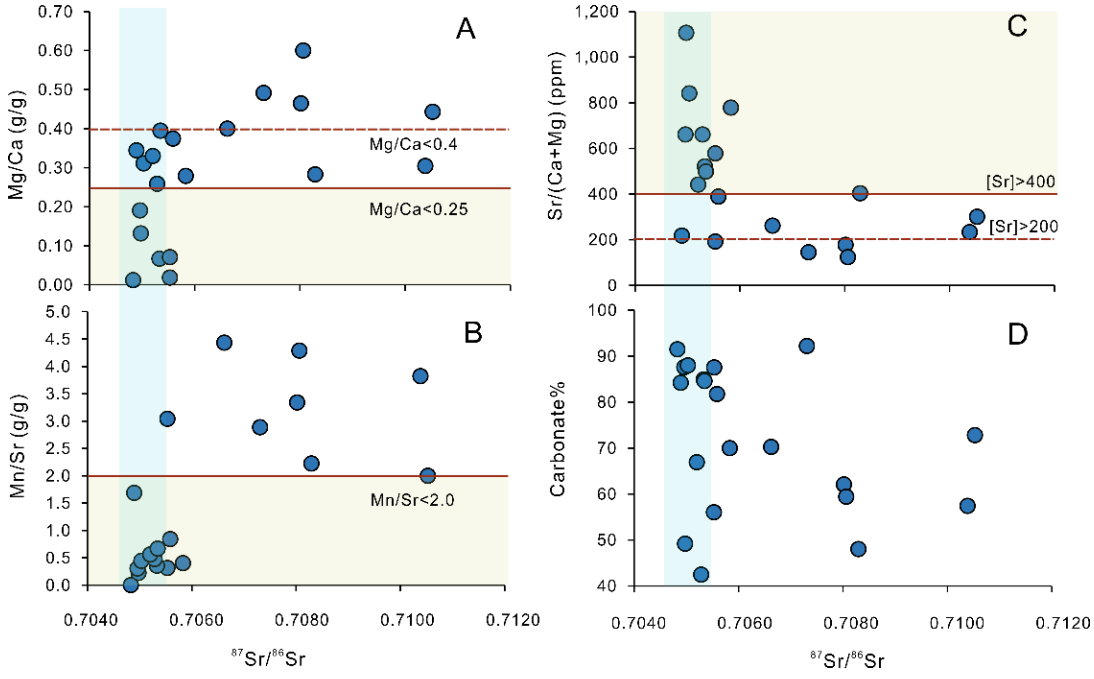

Fig. 11. Cross-plots bulk rock parameters of samples versus their lowermost Sr isotopes among three leaching cut-offs (see Table 4 for data). The blue shades show a seawater $^{87}Sr/^{86}Sr$ range (0.7046 ~ 0.7056) during this period. The yellow shades and red solid lines represent the thresholds highly probable for samples to yield seawater values. The red dash line shows recommended thresholds.

6. Conclusions

The major conclusions of this study are summarized below:

1) A Sr isotope leaching method for argillaceous and dolomitic limestones has been developed, whereby the most effective approach involves extracting the first 10% ~ 30% carbonate using weak acetic acid after ammonium acetate prewash.
2) REY patterns of carbonate rocks, in agreement with Sr isotopes, exhibit the most seawater-like patterns in the first leach after NH_4Ac prewash.
3) Organic matter remineralization during early diagenesis might influence the REY patterns of carbonate rocks but will not have much influence on Sr isotopes, while Sr isotopes are more vulnerable to dolomitization compared with REY. Both $^{87}Sr/^{86}Sr$ and REY in carbonates are sensitive to clay contamination.
4) Applying previously proposed leaching thresholds (e.g., preleach 30%~ 40%, 60%~70%) to the same argillaceous, dolomitic carbonate rocks leads to higher $^{87}Sr/^{86}Sr$ values compared with targeting the first 10% ~ 30% directly after NH_4Ac prewash. However, no significant differences were evident for high purity, and low Mg/Ca limestones, which underlines the importance of matching different sample types to the most appropriate dissolution method.

5) Thresholds for using Mg/Ca, [Sr] and Mn/Sr as screening tools for SIS study are recommended: Mg/Ca < 0.4 (preferably < 0.25) (g/g); [Sr] > 200 (preferably > 400) ppm; Mn/Sr < 2. Our leaching method increases the likelihood significantly of obtaining close to seawater Sr isotopic value for samples with carbonate content < 70% when other thresholds are met.

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Appendix A. Supplementary material

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