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Dear Editorial Manager,

I am enclosing herewith a manuscript entitled "*GEOCHEMICAL SURVEY AND POTENTIAL MINERALIZATION INVESTIGATION OF ILERO AND ITS ENVIRONS OYO STATE, NIGERIA.*"

This work reflects research on deciphering possible mineralization in the Ilero and its environments, Oyo State to its role in mineral exploration and exploitation.

I believe this submission will be useful to your readers, as it addresses some critical methodologies for discovering economic minerals. The research stands to benefit students researching Earth Science Studies.

This manuscript has been submitted to sedimentary geology journal. The current version is a preprint which has not yet been peer reviewed. Subsequent versions of this manuscript may have slightly different content. If accepted, the final version of this manuscript will be available via the 'Peer-reviewed Publication DOI' link on the right hand side of this webpage"

1 **GEOCHEMICAL SURVEY AND POTENTIAL MINERALIZATION INVESTIGATION OF ILERO AND**  
2 **ITS ENVIRONS OYO STATE, NIGERIA**

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6 Abstract

7 Ilero and its environs are part of the Kajola local Government area, Oyo State, and fall within the basement  
8 complex of southwestern Nigeria. The research work was focused on the geological mapping of the area and  
9 the study of the different rock and minerals units to determine the mineral resources of the area using  
10 petrographic, geochemical, and XRD analysis. The geochemical analysis showed the rocks and minerals to  
11 be diversely composed of 54-97% SiO<sub>2</sub>, 0.1-20% Al<sub>2</sub>O<sub>3</sub>, 0.8-4% MgO, 0.5-6% Fe<sub>2</sub>O<sub>3</sub>, < 1% K<sub>2</sub>O, TiO<sub>2</sub> and minor  
12 amounts of trace elements including copper, manganese, and lead. The mineralogy of the rocks and ore  
13 mineral comprised of quartz, biotite, hornblende, alkali feldspars, and tourmaline. There are also indications  
14 of the presence of chlorite and some ferromagnesian minerals such as pyroxene and olivine. The rocks are  
15 predominantly acid rocks along with a few basic rocks

16  
17 **Key word:** Petrographic, Provenance, Mineralization, Rocks, Ore Mineral.

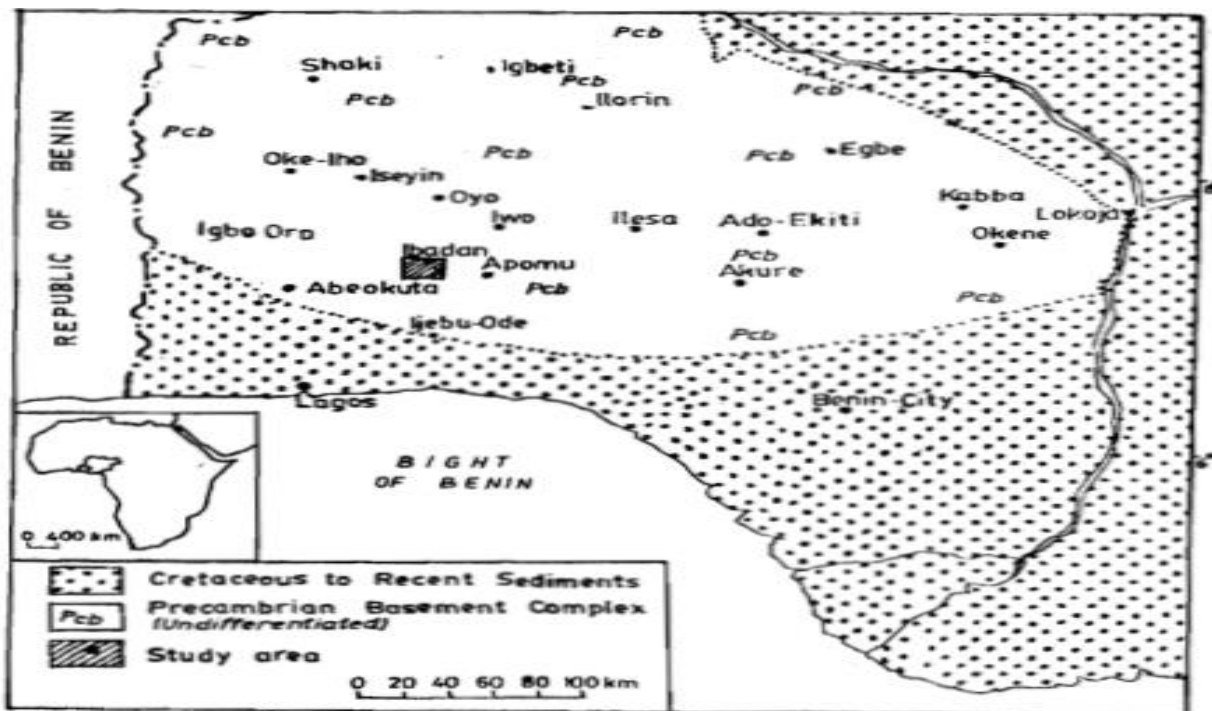
19 **1. Introduction**

20 The studied area is in the northwest of Okeho and Ilero, in Kajola local Government area, of Oyo-State.  
21 These areas fall within the Iseyin Schist belt, in the southwestern part of the Precambrian Complex geological  
22 province of Nigeria (Rahaman, 1976). As a result of less detailed work on the study area, not much  
23 information is available, and therefore the need for the delineation of various rock types of the different  
24 units of the basement complex in the area as well as the need for detailed petrographic and geochemical

25 studies on the rocks for assessment of their petrogenesis. The project work is aimed at revealing the  
26 mineralization potential of the Ilero area.

### 27 1.1 Geology and Location

28 The studied area is located around Ilero within, Kajola local Government area, Oyo State. It lies within  
29 longitude 003° 15'00" to 003° 26'00 and latitude 008° 05'00" to 008° 01' 00"N. The study area falls within  
30 the Precambrian Basement Complex of Southwestern Nigeria and in the Iseyin Schist Belt of southwestern  
31 Nigeria. It is characterized by occurrences of biotite granites, syenites, talcose, amphibolites, and bodies of  
32 pegmatites and quartz-veins. The study area is accessible from Oyo-Iseyin, Okeho- Ilero, Igana, and Shaki -  
33 Ipapo -Otu Road (Rahaman, 1976). Numerous footpaths and un-tarred rural roads were also accessible. The  
34 study area is more accessible during the dry season when the field aspect of this study was carried out.  
35 Accessibility would have been difficult during the rainy season when the drainage channels would have been  
36 overflowed and the forest will be thick thereby causing coverage of exposures and making navigation within  
37 the area difficult.



38

39 Fig. 2 Map of Regional Geology of the Study Area (Oyo State) within the basement complex in  
40 southwestern Nigeria (Rahaman, 1976).  
41

## 42 **2. Materials and Methods**

### 43 *2.1 The Petrographic Analysis*

44 The petrographic analysis was used to study the mineralogical composition of rock samples by  
45 studying the behavior of these minerals under the plane and crossed polarized light (Ekeleme, 2017).  
46 This study involves cutting a thin section of rock samples and critically studying these thin sections  
47 under a polarizing microscope to determine the optical properties of the samples. Basically, the thin  
48 section is made of a glass slide, a thin film of a rockslide (30 microns), and a cover slip. The minerals  
49 used in the preparation of thin sections include Canada Balsam, Carborundum powder of 400, 600,  
50 and 800 grits, and Araldite. The equipment used was the Logitech rock cutting machine, hot plate,  
51 smooth glass, glass slide, diamond pen, etc. (Nwachukwu, 2011). The rock samples to be studied were  
52 first cut into slices of thickness 3mm -4mm using the thin section (Logitech) machine, then the surface  
53 of the rock was then smoothed by continuously grinding its surface on a glass plate using coarse grit  
54 carborundum (Folorunso, 2013). The planned surface of the rock sample is then mounted on a glass  
55 slide with the use of Araldite and pressed together to remove air bubbles using forceps and placed on  
56 the hot plate for a few minutes. The other end of the rock sample is further lowered by cutting a certain  
57 thickness off using the cutting machine and is then ground on the glass plate using coarse grit  
58 carborundum powder followed by fine grit carborundum powder until the desired thickness of 30  
59 microns (0.03mm) is achieved. At this thickness, light can pass through the rock sample thereby  
60 revealing the mineral composition of such rock sample. A cover slip is used to cover the section with  
61 the aid of the Canada balsam to prevent the composition from being altered. After the thin section of  
62 the rock, samples have been successfully prepared, then the mineralogical composition of these rock  
63 samples is observed under a polarizing microscope.

### 64 *2.2 Geochemical Analysis*

65 The first step in carrying out a geochemical analysis is pulverization. After the rock samples are  
66 selected for geochemical analysis, they are broken into small fragments with each fragment obtained  
67 from the cleanest, freshest portion of the entire rock sample to obtain distinct compositional results  
68 of the samples (Black, 1980). These small rock fragments are then pulverized (ground) such that there  
69 are no conspicuous grains observed in the sample. It is important to immediately transfer the  
70 pulverized sample into a tight sample holder to protect the composition of the bulk (Caby, 1981). The

71 geochemical analysis is basically used to analyze rock geochemistry, especially for the major, minor, and  
72 trace elements present in the rock and ore samples (Cox, 1979). The essence of geochemistry is to  
73 relate mineralogy with the chemistry of the rock. The major, minor, and trace element geochemical  
74 compositions were obtained by X-Ray fluorescence (XRF) while X-Ray diffraction (XRD) was used for  
75 mineralogical composition (Dada, 2006).

### 76 2.3 X-Ray Diffraction (XRD)

77 X-Ray diffraction is one of the important techniques for material characterization. The diffraction  
78 experiments using x-rays help to study the structural properties of materials on an atomic scale  
79 (Ahmed, 2015). The techniques are also used to measure crystallite size and calculate lattice strain,  
80 and chemical composition, and determine phase diagrams as well (Arif, 2017). They are also one of the  
81 most useful in the characterization of crystalline materials, such as metals, inter-metallic, ceramics,  
82 minerals, polymers, plastics, or other inorganic or organic compounds. X-ray diffraction techniques  
83 can be used to identify the phases present in samples from raw starting materials to finished product  
84 and to provide information on the physical state of the sample, such as grain size, texture, and the  
85 crystal perfection. Most x-ray diffraction techniques are rapid and non-destructive, some instruments  
86 are portable and can be transported to the sample (Ahmed, 2015). The type of information obtained  
87 from x-ray diffraction studies ranges from sample composition to details of the crystal structure or the  
88 state of orientation of the crystallites. Phase identification can be conducted on virtually all single  
89 crystal or powder samples (Arif, 2017). Also, useful are measurements of the physical state of a sample  
90 that detect differences from the ideal crystal (presence of defects, strain, etc). In powder diffraction,  
91 finely ground and homogenized samples are used. The sample may be held in thin-walled glass or  
92 cellophane capillary tubes (Zandomeneghi, 2010). The x-rays (usually, K-radiation of wavelength  
93 1.5418Å emitted from copper) required for diffraction are mono-chromatized by filtering with the help  
94 of foils or crystal monochromators. These x-rays are collimated and directed onto samples  
95 (Rahman, 2016). In powder form, a significant number of crystallites can be expected to be oriented  
96 to fulfill the Bragg condition for reflection from every possible interplanar spacing.

97

98

99 **3. Results and Discussion**

100 *3.1 Geochemical*

101 The geochemical analysis showed the rocks and minerals to be diversely composed of 54-97% SiO<sub>2</sub>,  
102 0.1-20% Al<sub>2</sub>O<sub>3</sub>, 0.8-4% MgO, 0.5-6% Fe<sub>2</sub>O<sub>3</sub>, < 1% K<sub>2</sub>O, TiO<sub>2</sub> and minor amounts of trace elements  
103 including copper, manganese, and lead. These geochemical concentrations values obtained within the  
104 study area are close to that reported by (Danbatta, 2007) for Precambrian rocks/ minerals in the Zuru  
105 Schist Belt characterized by extremely low K<sub>2</sub>O; high Al<sub>2</sub>O<sub>3</sub>; and low TiO<sub>2</sub>, Cs, Ge, and Rb contents which  
106 were in accordance with the orogenic nature of the Pan-African event.

107 The results of the geochemical analysis show that altered pegmatite (SO-1), syenite (SO-2), and  
108 quartz vein (SO-4) are constituted with 54.8%, 63.61%, 98.86%, SiO<sub>2</sub>, low concentrations of  
109 K<sub>2</sub>O, CaO, Fe<sub>2</sub>O<sub>3</sub>, MgO, TiO<sub>2</sub>, K<sub>2</sub>O, Na<sub>2</sub>O and P<sub>2</sub>O<sub>5</sub> and minor trace elements (Table 1 and Fig. 3a-e).  
110 These geochemical composition values suggest that the rocks underlying the area of Ilero and Okeho  
111 possess a high abundance of silica that would support the granitic origin of the rocks (Dada, 2006). The  
112 amphibolite (SO-3) is however significantly rich in iron oxide, magnesia, and calcium oxide as to be  
113 appreciated to be basic lithology. The Tourmaline ore mineral (SO-5) is a ubiquitous accessory mineral  
114 in rocks of the Earth's crust. They can adjust their composition to suit a wide variety of environments,  
115 and therefore display a remarkable range of instability in terms of pressure, temperature, fluid  
116 composition, and host-rock composition.

117 The results of the geochemical analysis show that altered pegmatite (SO-1), syenite (SO-2), and quartz  
118 vein (SO4) are constituted with 54.8%, 63.61%, 98.86%, SiO<sub>2</sub>, low concentrations of K<sub>2</sub>O, CaO, Fe<sub>2</sub>O<sub>3</sub>,  
119 MgO, TiO<sub>2</sub>, K<sub>2</sub>O, Na<sub>2</sub>O and P<sub>2</sub>O<sub>5</sub> and minor trace elements (Table 4.1 and Fig. 4.1-4.5). These  
120 geochemical composition values suggest that the rocks underlying the area of Ilero and Okeho possess  
121 a high abundance of silica that would support the granitic origin of the rocks (Dada, 2006). The  
122 amphibolite (SO-3) is however significantly rich in iron oxide, magnesia, and calcium oxide as to be  
123 appreciated to be basic lithology. The Tourmaline ore mineral (SO-5) is a ubiquitous accessory mineral  
124 in rocks of the Earth's crust. They can adjust their composition to suit a wide variety of environments,  
125 and therefore display a remarkable range of stability in terms of pressure, temperature, fluid

126 composition, and host-rock composition The correlation matrix for the data obtained (Table 2) shows  
 127 a wide variation in the correlation coefficient (r) between element pairs. For example, the r values  
 128 range from -0.05 between Rb and Al to 0.99 between Rb and Pb. A strong to very strong positive  
 129 correlation occurs among Ti, Fe, S, and Cu positively correlated with each other. Considerable high  
 130 correlation was found between Rb and Pb (R = 0.99), S and Al (R = 0.69), S and Ti (R=0.81), S and Cu  
 131 (R=0.80), Cu and Al (R=0.80), Ti and Fe (R = 0.78). Other elements also showed a fair to strong  
 132 correlation as summarized in Table 2, (Okonkwo, 2013). According to Okonkwo and Folorunsho  
 133 (2013), the Fairly strong to very strong positive correlation between Ti, Fe, S, Cu, Al, Rb, and Pb in the  
 134 analyzed Rocks suggests the possibility of Ti-Fe-S-Cu-Al- Rb-Pb bearing minerals within the rocks of the  
 135 study area.

136 Table 1. Result of Geochemical Analysis of Rocks and Minerals from Ilero Area.

	SO-1	SO-2	SO-3	SO-4	SO-5
Rock/Mineral	Altered Pegmatite	Syenite	Amphibolite	Vein Quartz	Tourmaline (Ore mineral)
SiO <sub>2</sub>	54.80	63.61	39.23	96.86	61.82
TiO <sub>2</sub>	0.05	0.70	0.97	0.02	0.59
Al <sub>2</sub> O <sub>3</sub>	20.67	14.99	21.16	0.10	12.15
Fe <sub>2</sub> O <sub>3</sub>	0.57	5.19	30.98	0.54	3.72
MnO	0.02	0.32	1.32	0.03	0.37
MgO	4.20	3.82	3.69	0.79	5.85
CaO	0.45	0.44	4.38	0.04	3.66
Na <sub>2</sub> O	3.36	1.50	0.05	0.00	2.00
K <sub>2</sub> O	9.00	3.08	0.10	0.07	4.59
P <sub>2</sub> O <sub>5</sub>	0.12	0.15	3.61	0.31	0.50
LOI	22.12	3.00	2.60	0.60	2.12
SO <sub>2</sub>	0.04	0.05	0.08	0.00	0.08
CuO	0.03	0.01	0.03	0.00	0.03
PbO	0.01	0.02	0.00	0.07	0.01
Rb <sub>2</sub> O	0.01	0.02	0.00	0.00	0.01
Cs <sub>2</sub> O	0.02	0.00	0.02	0.03	0.00
GeO <sub>2</sub>	3.08	1.00	1.03	0.40	1.00

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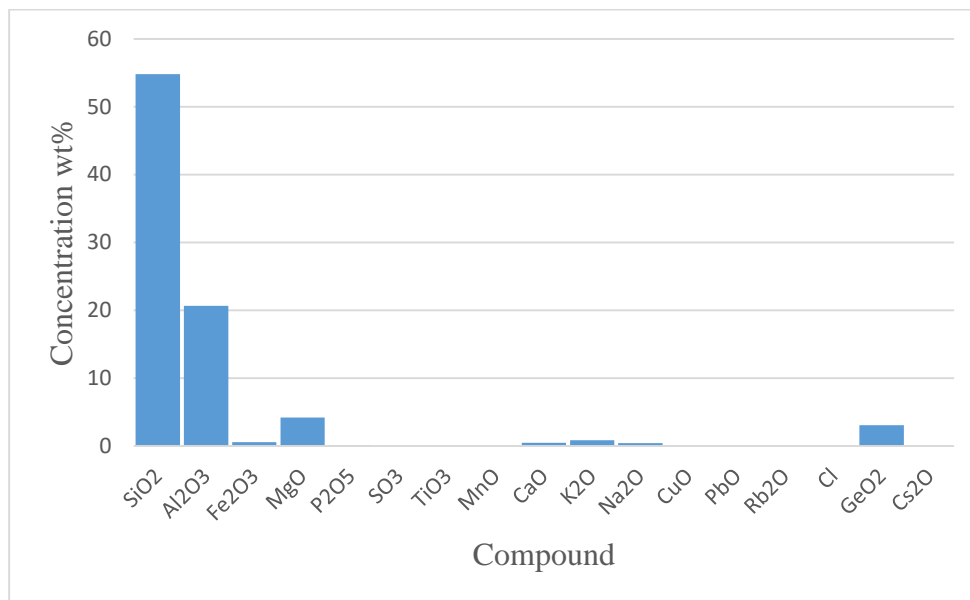
141 Table 2: Correlation between Ore Minerals of the Study Area

	Al	Fe	S	Ti	Cu	Pb	Rb	Ge
Al	1							
Fe	0.50	1						
S	0.69	0.58	1					
Ti	0.49	0.78	0.81	1				
Cu	0.80	0.39	0.79	0.36	1			
Pb	-0.16	-0.36	0.40	0.10	0.25	1		
Rb	-0.05	-0.26	0.51	0.18	0.37	0.99	1	
Ge	0.64	-0.19	0.07	-0.33	0.56	-0.16	-0.10	1

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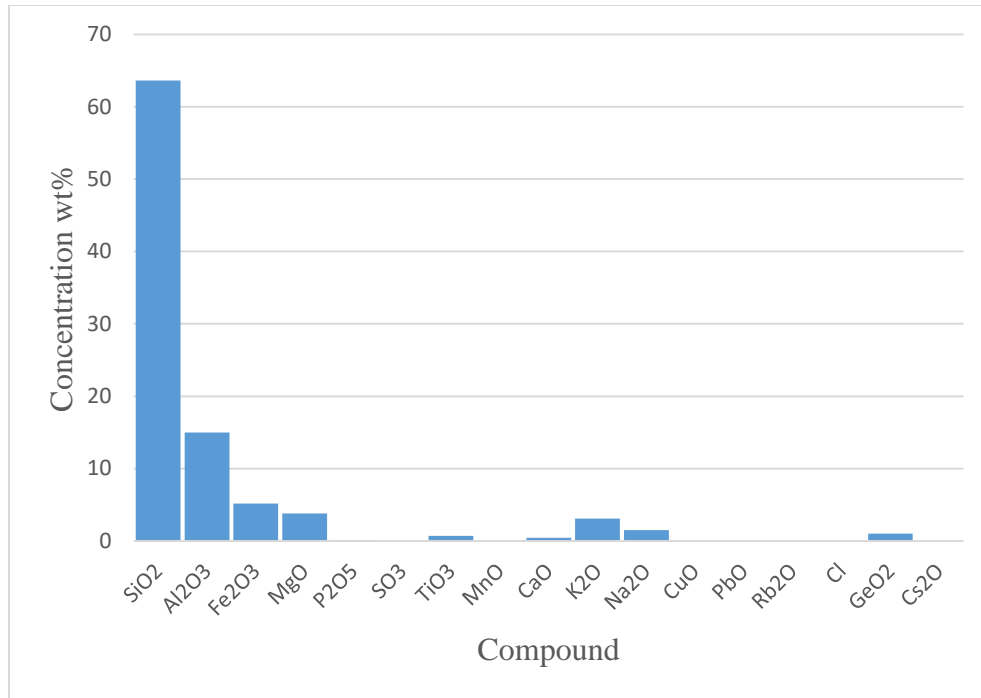
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Fig 3a. Distribution of Major Oxides in Tourmaline (Ore Mineral).

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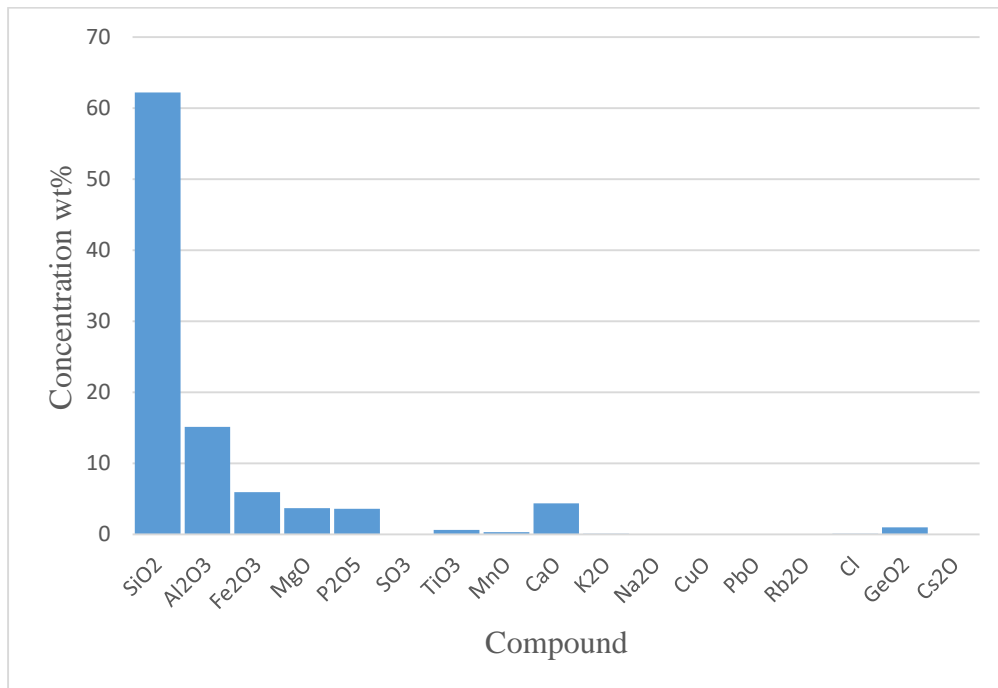


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Fig 3b. Distribution of Major Oxides in Altered Pegmatite.

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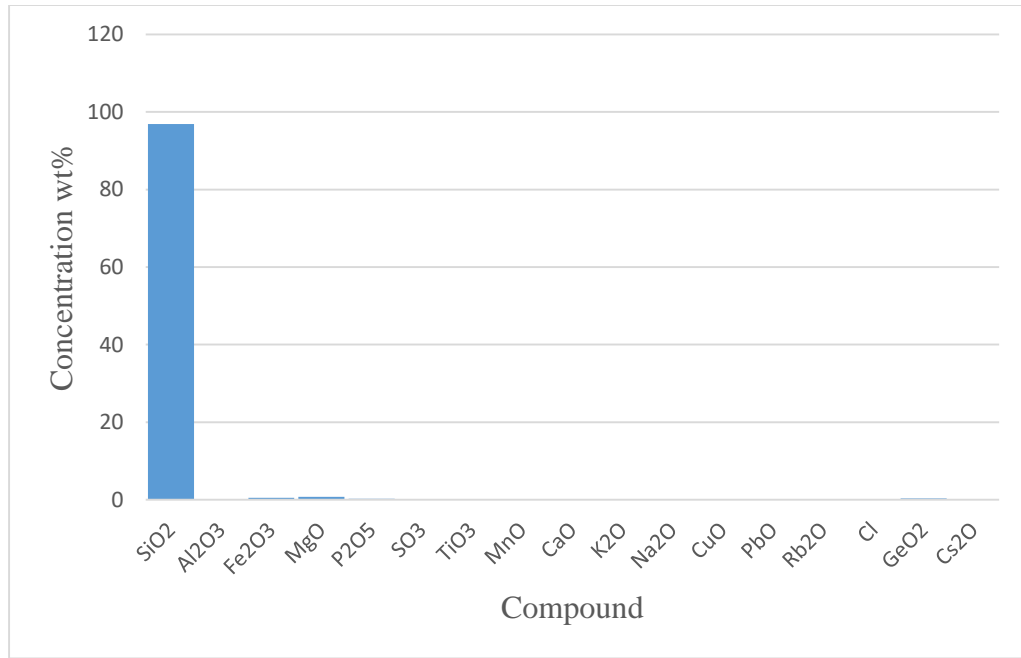


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Fig 3c. Distribution of Major Oxides in Syenite.

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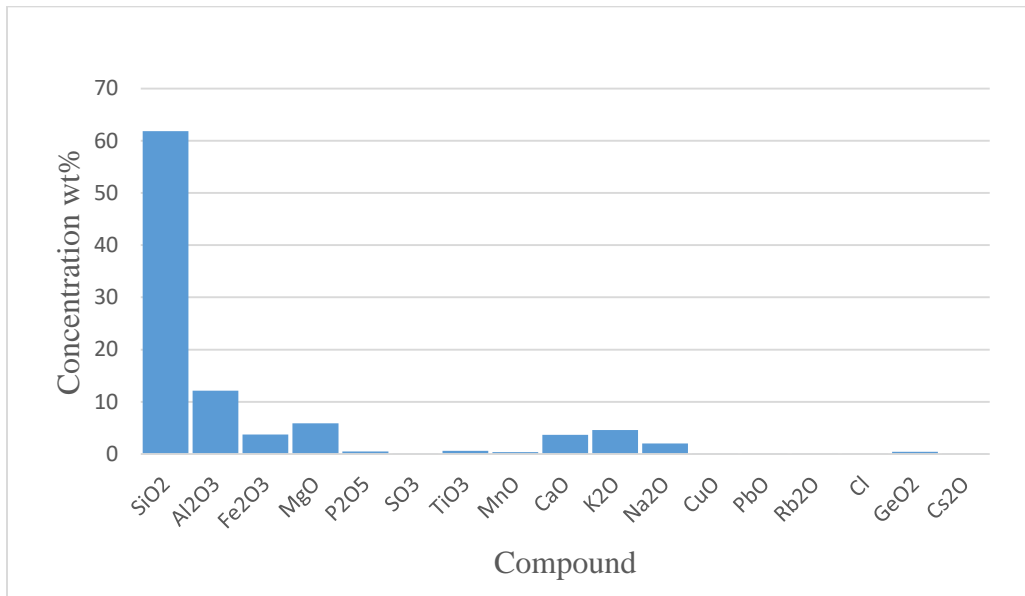
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Fig 3d. Distribution of Major Oxides in Vain Quartz



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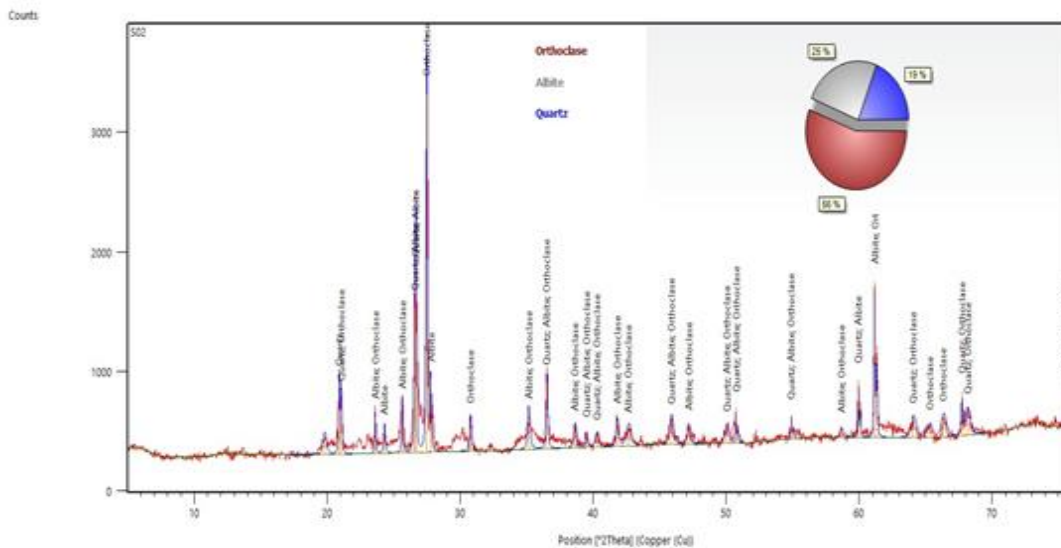
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Fig 4e. Distribution of Major Oxides in Amphibolite

162 3.2 X-Ray Diffraction Analysis (XRD)

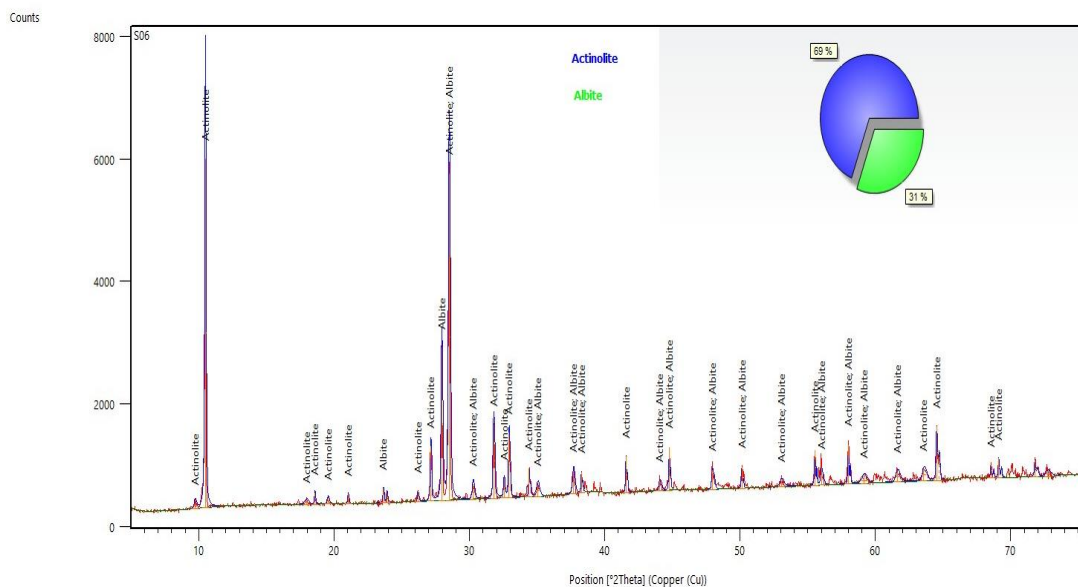
163 The x-ray diffraction results are presented as charts in Figs. 4a-c. The mineral phases were  
164 identified with respect to the International Committee on Powder Diffraction File (ICPDF) and their  
165 abundances computer calculated (Zandomeneghi, 2010).

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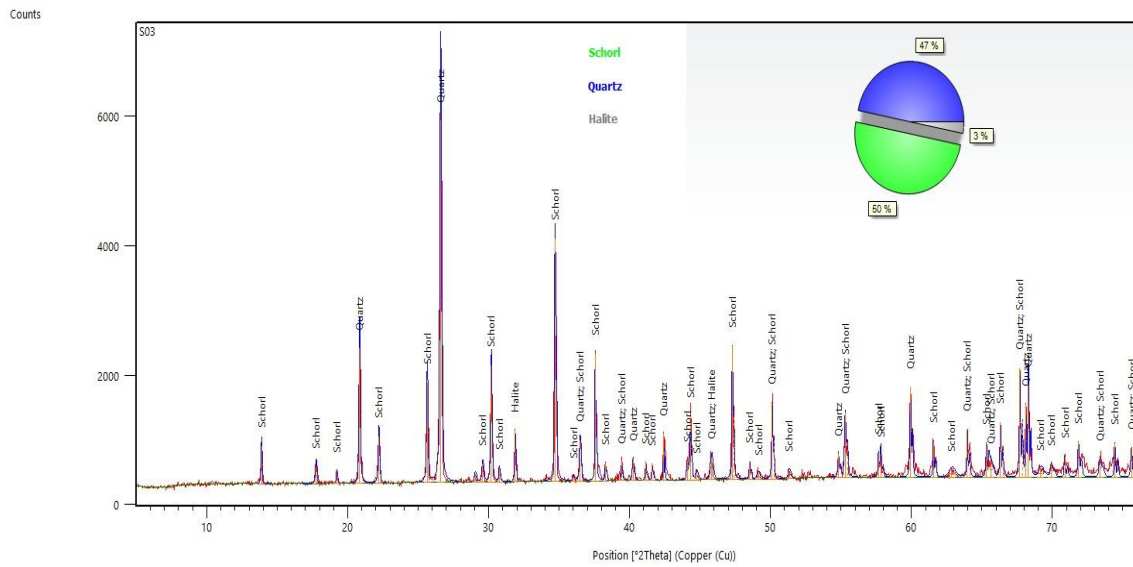
168 Fig. 4a. XRD result of Syenite.



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170 Fig. 4b XRD Chart of Amphibolite.

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Fig. 4c. XRD result of Tourmaline (Ore mineral)

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175 Table 3. Minerals Identified (Names of compounds and formulae) in Sample by X-Ray Diffraction

Sample I.D	Mineral Identified	Chemical Formula	Crystal System	Percentage (%)
SO-2	Orthoclase	$KAlSi_3O_8$	Tectosilicate	56%
	Albite	$NaAlSi_3O_8$	Tectosilicate	25%
	Quartz	$SiO_2$	Hexagonal	19%
SO-3	Schorl	$NaFe^{3+}Al_6(BO_3)_3Si_6O_{18}(OH)_4$	Prismatic	50%
	Quartz	$SiO_2$	Hexagonal	47%
SO-6	Actinolite	$Ca_2(Mg_5Fe^{2+})Si_8O_{22}(OH)_2$	Monoclinic	69%

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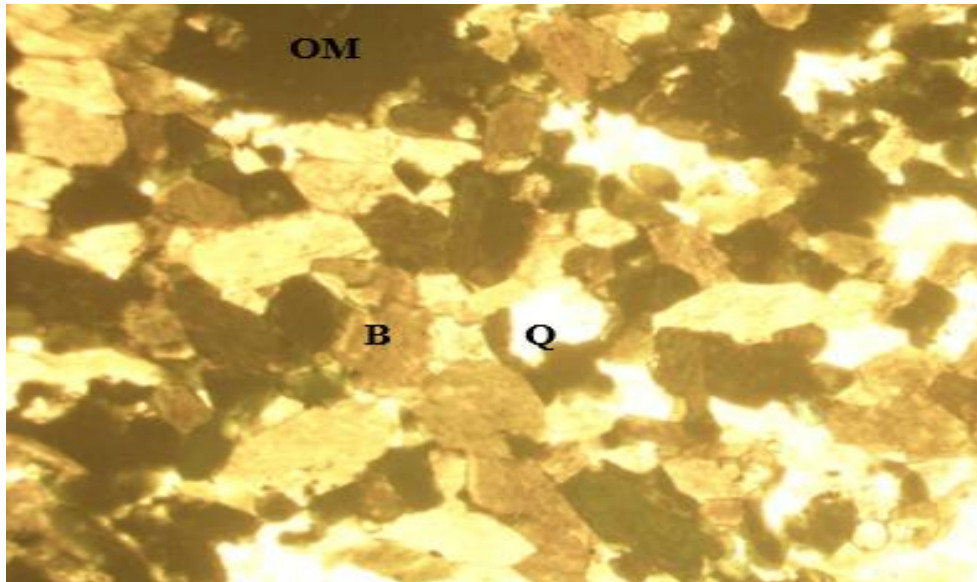
178 (Table 3) showed the XRD results of the mineral composition of the rocks within the Ilero area.  
179 The XRD results showed that Albite ( $\text{NaAlSi}_3\text{O}_8$ ) is found except for Tourmaline (Rahman, 2017). The  
180 Syenite apart from being rich in Orthoclase had quartz ( $\text{SiO}_2$ ) and Albite. Syenite is a coarse-grained  
181 intrusive igneous rock with a general composition like that of granite, but deficient in quartz, which, if  
182 present at all, occurs in relatively small concentrations (< 5%). Amphibolite had Actinolite and Albite.  
183 Amphibolite is a dark green foliated rock predominantly composed of green hornblende, plagioclase,  
184 and quartz, and a minor occurrence of black pyroxene, quartz, and epidote. Tourmaline (ore mineral)  
185 showed Schorl, and Quartz (Rahman *et al.*, 2017; Arif, 2017). Tourmaline generally occurs as a minor,  
186 accessory phase in rocks, but tourmalines, where tourmaline is a major mineral, are well known (Arif,  
187 2017). Tourmaline forms in a wide variety of settings, including as a diagenetic mineral in buried  
188 sedimentary basins, as a gangue mineral in ore deposits, associated with a contact, regional and  
189 subduction-related metamorphism, in metasomatism, and as a significant mineral typically crystallizing  
190 in fractionated igneous bodies (Arif, 2017). The peaks on the diffraction spectra (Fig. 4.6-4.8) signified  
191 the abundance of each of these minerals in their different samples.

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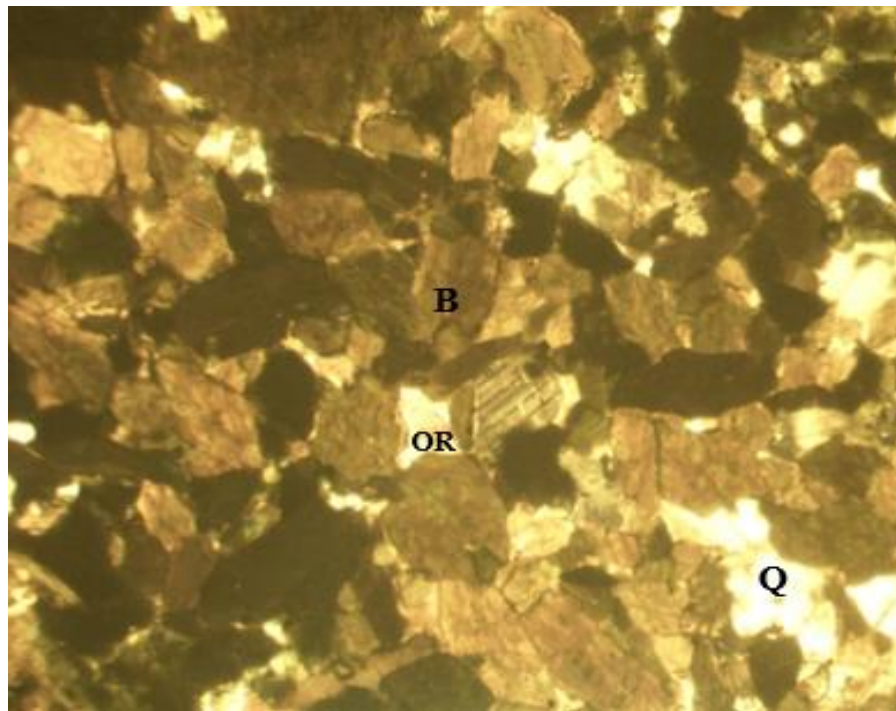
### 193 3.3 Petrology and Heavy Mineral Study

194 Thin sections of rocks and ore minerals (Tourmaline (Ore mineral) from Ilero environs studied  
195 under the petrological microscope showed the rocks (Altered Pegmatite, Syenite, Amphibolite, and  
196 Vein Quartz) to be coarse-grained textured with euhedral mineral crystals (Nwachukwu, 2011). In  
197 general, the minerals present include quartz (20%), orthoclase (6%), hornblende (7%), tourmaline (2%),  
198 and biotite (40%) (Black, 1980). (Fig 5a-j) shows the minerals observed both in plain polarized light  
199 (PPL) and under crossed Nicol (XPL). Quartz and biotite are the most abundant minerals observed.  
200 Quartz constitutes about 20% of the mineral content of the Vein Quartz and its morphology rarely  
201 forms euhedral crystals except in some veins which are anhedral (Danbatta, 2007). Quartz is an  
202 extremely common mineral in continental crustal rocks such as Granite, Pegmatite, etc. It is the most  
203 common vein-forming mineral. Biotite is one of the most common rock-forming minerals (Folorunso,  
204 2013). It is a major constituent of rocks of intermediate composition, such as granodiorites, diorites,  
205 monzonites, and some nepheline syenites.

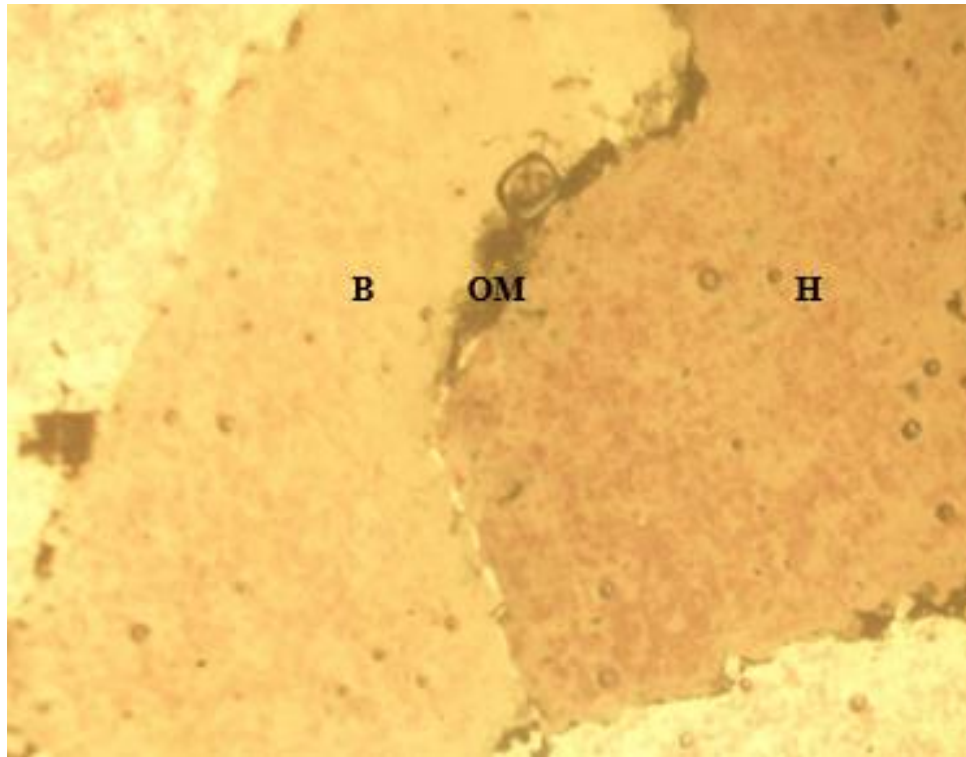
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207  
208 Fig 5a: Photomicrograph of sample Vein Quartz. showing Q-Quartz, B- Biotite, OM-Opaque Mineral  
209 Magnification x4.  
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212 Fig 5b: Photomicrograph of sample Altered Pegmatite. showing Q-Quartz, B-Biotite, OR-Orthoclase Feldspar  
213 Magnificationx4.

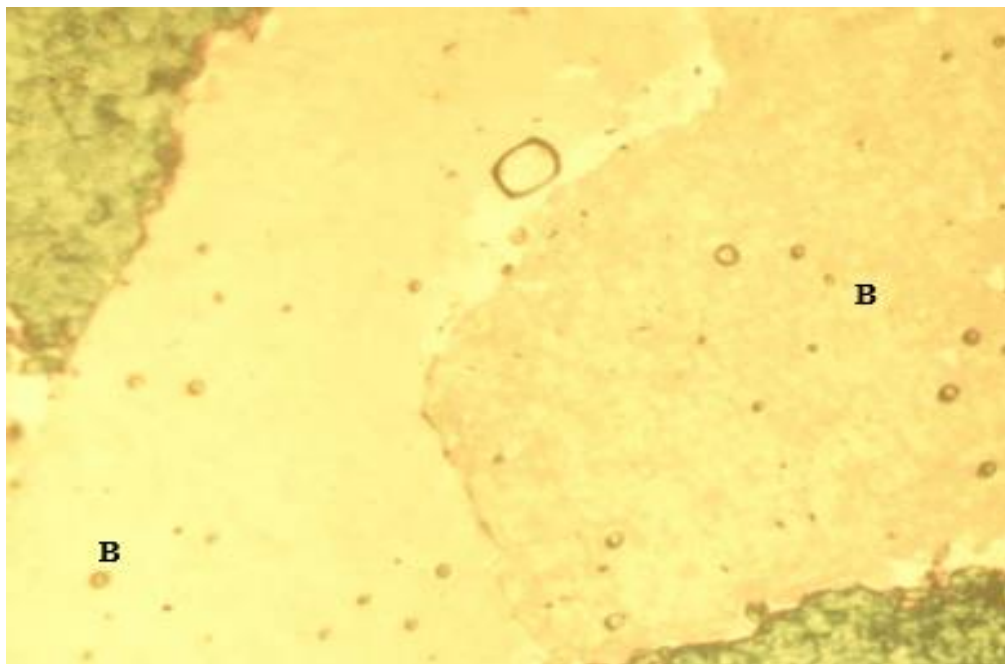


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Fig 5c: Photomicrograph of Amphibolite showing B-Biotite, H-Hornblende. Magnification x4

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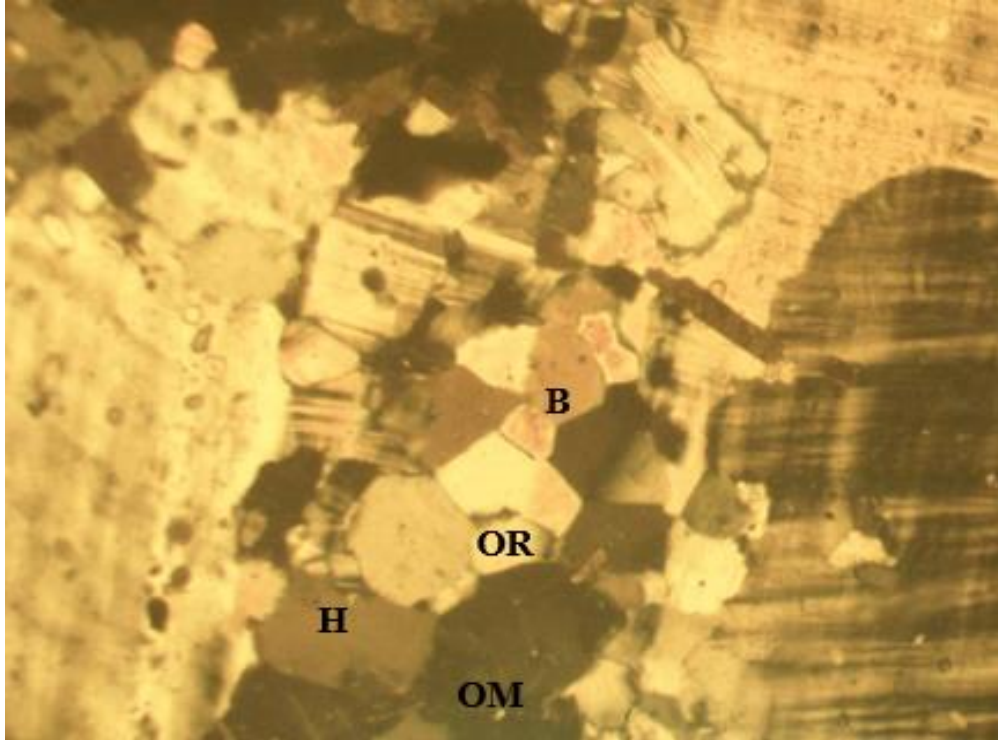


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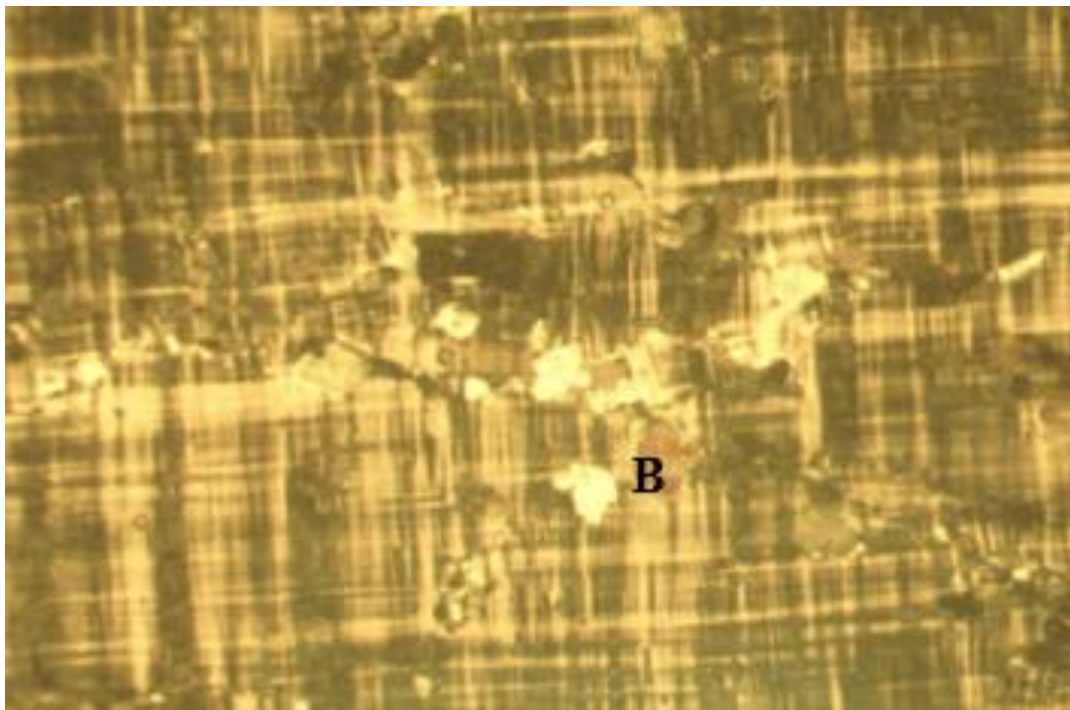
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Fig 5d: Photomicrograph of Vein Quartz, showing B-Biotite. Magnification x4

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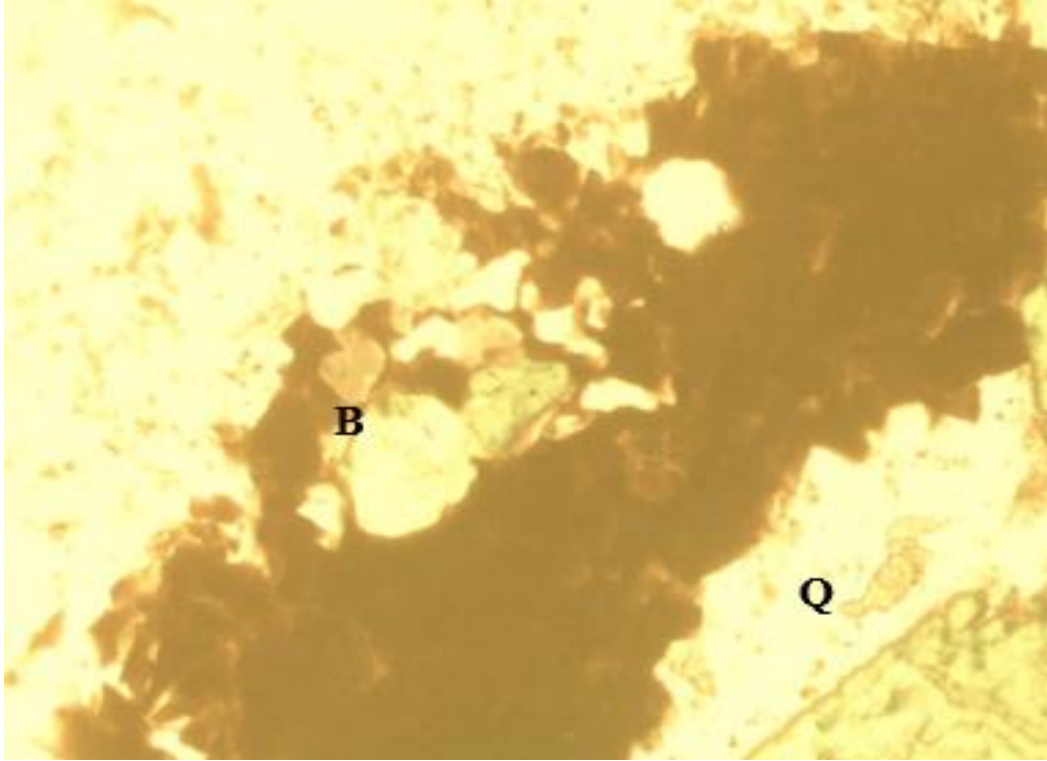


220  
221 Fig 5e: Photomicrograph of Amphibolite showing B- Biotite, OR-Orthoclase Feldspar, H- Hornblende, OM-  
222 Opaque Mineral. Magnification



223  
224 Fig 5f: Photomicrograph of sample Vein Quartz showing B-Biotite Magnification x4  
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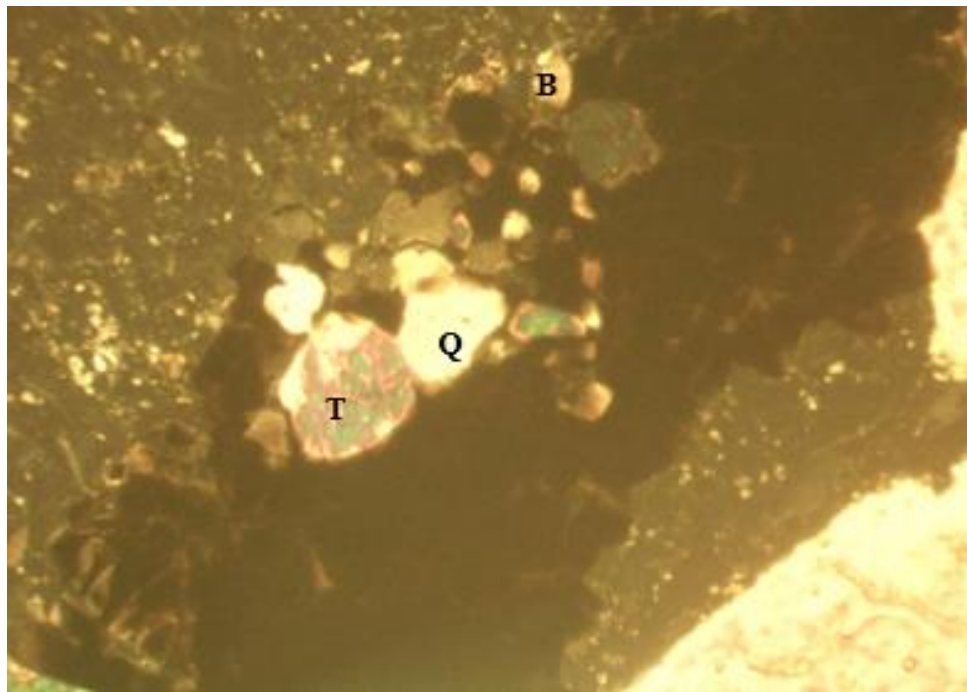


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Fig5g: Photomicrograph of Vein Quartz showing B-Biotite, Q- Quartz Magnification x4



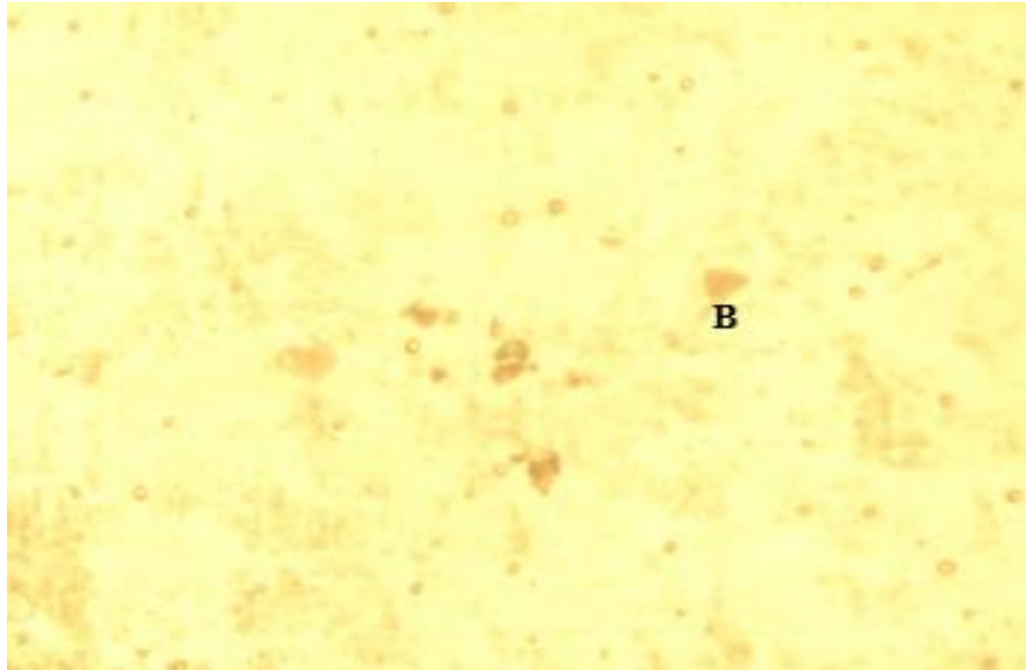
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Fig 5h: Photomicrograph of Tourmaline (Ore Mineral) showing B-Biotit, T-Tourmaline, Q-Quartz. Magnification x4

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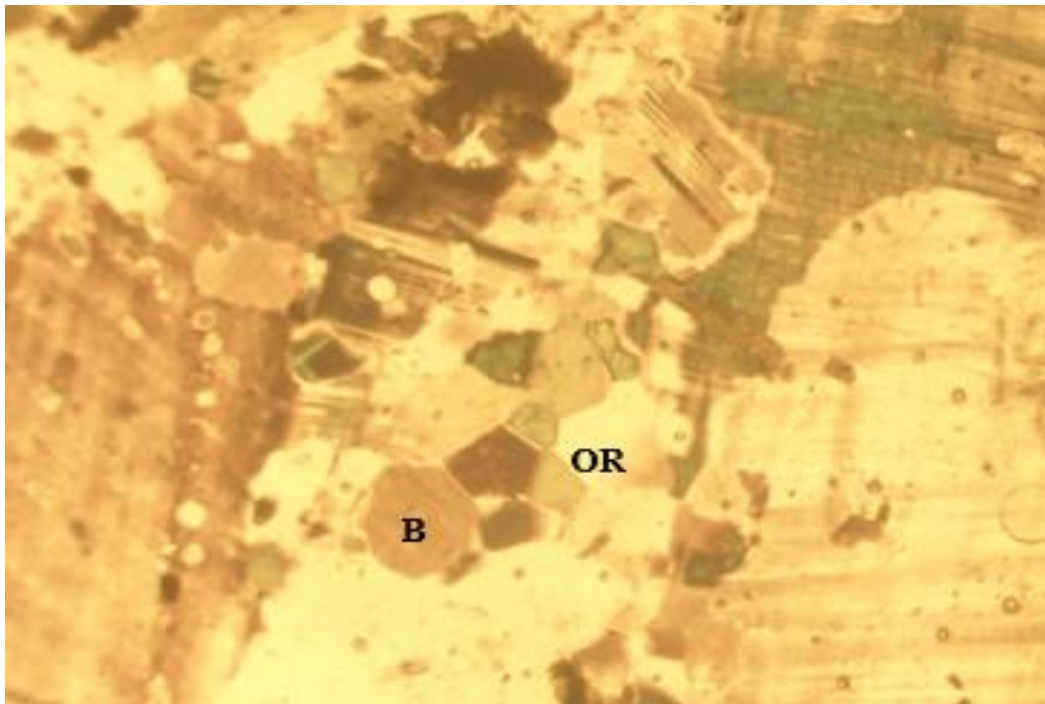


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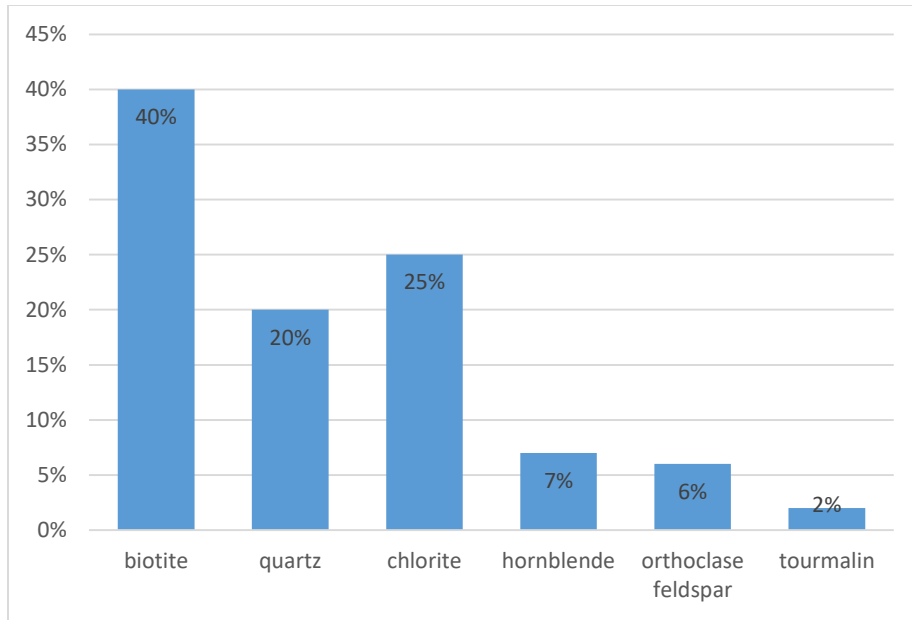
Fig 5i: Photomicrograph of Vein Quartz showing B-Biotite. Magnification x4

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237 Fig 5j: Photomicrograph of sample Altered Pegmatite showing B-Biotite, OR-Orthoclase Feldspar. Magnification  
238 x4



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240 Fig 5: Bar chart showing estimated modal mineralogical composition of granite under the microscope.  
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251 Table 4.: Petrographical Analysis of Ilero Samples

Minerals	Optical Properties	Crossed Polarized (xpl)	Plane Polarized (ppl)	Percentage in Thin Section
Quartz	Color	Bluish to milky	Colourless	20%
	Relief:	Low relief		
	Form/size	Anhedral		
	Inclusion	Absent		
	Cleavage	Absent		
Biotite	Color	Brown	Brown	40%
	Relief:	Low		
	Form/size	Euhedral		
	Inclusion	Absent		
	Cleavage	Two (2) cleavage direction		
Chlorite	Color	Green	Green	25%
	Relief:	Low		
	Form/size	Euhedral		
	Inclusion	Absent		
	Cleavage	Present		
Hornblende	Color	Brown	Brown	7%
	Relief:	High		
	Form/size	Prismatic		
	Inclusion	Present		
	Cleavage	Present		
Tourmaline	Color	Dark Brown /Yellow	Dark Brown	2%
	Relief:	High		
	Form/size	Slender Prism		
	Inclusion	Absent		
	Cleavage	Present		
Orthoclase Feldspar	Color	Grey	Grey	6%
	Relief:	Low		
	Form/size	Prismatic		
	Inclusion	Present		
	Cleavage	Present		
Total				100%

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256 **3. Conclusion**

257 The correlation analysis of some selected elements from the result of the geochemical analysis shows  
258 Ti, Fe, S, Cu, Al, Rb, and Pb among others to be strongly correlated. The conclusion drawn from the  
259 integration of the geochemical analysis, and correlation analysis in this study is that the study area is  
260 mineralized with gibbsite-cuprite-pyrite-galena-cassiterite-pollucite. The result of the petrographic study  
261 reveals that the rocks and ore minerals in the study area are predominantly whitish-grey-greenish in color  
262 with the textural compositions of porphyritic to fine-grained structures. Petrographically, the rocks and ore  
263 minerals in the study area are composed of major rock-forming silicate minerals such as quartz, biotite,  
264 hornblende, tourmaline, and feldspars. The XRD results showed that Albite, Orthoclase, quartz, Schorl, and  
265 Actinolite are possible mineralized minerals in the samples of the study area. Based on the petrographical,  
266 XRD, and geochemical analysis obtained, the results suggest that the study area possesses highly  
267 mineralized with economical minerals.

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269 REFERENCES

- 270 Ahmed, Z.A., Maxim, L., Sarah, J.V., Michael, L.J., Ahmed, B. 2015. [Pore-scale analysis of formation damage](#)  
271 [in Bentheimer sandstone with in-situ NMR and micro-computed tomography experiments](#), [Journal](#)  
272 [of Petroleum Science and Engineering](#), v. 129. P. 48-57.
- 273 Arif, M., Ahmed, B., Maxim, L., Stefan, I. 2017. [Impact of solid surface energy on the stability of](#)  
274 [CO<sub>2</sub>/brine/mineral systems as a function of pressure, temperature, and salinity](#), [Energy](#)  
275 [Procedia](#),114, p. 4832-4842.
- 276 Black, R .1980. [Precambrian of West Africa](#). [EPISODES Int J Geosci](#). 1980; 4:38.
- 277 Dada, S. S. 2006. [Proterozoic evolution of Nigeria: In Oshi The basement complex of Nigeria and its mineral](#)  
278 [resources \(A tribute to prof MAO Rahaman\)](#), [akin jihad and company Ibadan](#). [J Sci Res](#). 2006;29-44.
- 279 Danbatta U.A. 2007. [Geochemistry and petrogenesis of Precambrian amphibolites in the Zuru schist belt,](#)  
280 [northwest Nigeria](#). [J Min Geol](#). 2007; 43:23-30.

- 281 Ekeleme I.A, Uzoegbu M.U, Abalaka I.E. 2017. Petrographic evaluation of rocks around Arikya and its  
282 environs, North Central Nigeria. *Int J Geol Min.* 2017; 3:103-109.
- 283 Folorunso A.F, Ayolabi E.A. 2013. Geological mapping, petrological study, and structural analysis of  
284 Precambrian basement complex rocks in part of Ago-Iwoye Southwestern Nigeria. *Int Res J Geol*  
285 *Min.* 2013; 3:19-30.
- 286 Nwachukwu M.A. 2011. Petrographic analysis for naming and classifying an igneous intrusive rock in the  
287 Lower Benue Trough. *J Geol Min Res.* 2011; 3:63-72.
- 288 Okonkwo, C. T., and Folorunso, I. O. 2013. Petrochemistry and geotectonic setting of granitic rocks in  
289 Aderan area, SW Nigeria. *Journal of Geography and Geology*, vol. 5(1), pp. 30-44.
- 290 Opara K.D, Obioha Y.E, Onyekuru S.O, Okereke C. 2014. Petrology and geochemistry of basement complex  
291 rocks in okom-ita area, Oban Massif, southeastern Nigeria. *Int J Geosci.* 2014; 5:394-407.
- 292 Rahaman, M.A. 1976. Review of the basement geology of southwestern Nigeria. In. *Geology of*  
293 *southwestern Nigeria In Geology of Nig (C.A Kogbe ed) Elizabethan Publishing Co. Lagos.,pp: 41-56.*
- 294 Rahman, T., Maxim, L., Ahmed, B., Stefan, I. 2016. Residual trapping of supercritical CO<sub>2</sub> in oil-wet  
295 sandstone, *Journal of Colloid and Interface*, v. 469, p. 63-68.
- 296 Rahman, T., Maxim, L., Yihuai, Z., Ahmed, B., Stefan, I. 2017. Influence of rock microstructure on its electrical  
297 properties: an analysis using X-ray microcomputed tomography, *Energy Procedia*, v. 114, p. 5023-  
298 5031.
- 299 Zandomeneghi, D., Voltolini, M., Mancini, L., Brun, F., Dreossi, D., Polacci, M. 2010. Quantitative analysis  
300 of x-ray microtomography images of geomaterials: Application to volcanic rocks, *Geosphere*, v. 6,  
301 no. 6, p.793-804.

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