

1 **Persistent effect of El Niño on global economic growth**

2 Christopher W. Callahan^{1,2*} & Justin S. Mankin^{2,3,4}

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4 ¹Program in Ecology, Evolution, Environment and Society, Dartmouth College, Hanover, NH

5 ²Department of Geography, Dartmouth College, Hanover, NH

6 ³Department of Earth Sciences, Dartmouth College, Hanover, NH

7 ⁴Ocean and Climate Physics, Lamont-Doherty Earth Observatory of Columbia University, Palisades, NY

8 *Corresponding author, Christopher.W.Callahan.GR@dartmouth.edu

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14 *corresponding author with questions or comments at Christopher.W.Callahan.GR@dartmouth.edu or*
15 *@cwwcallahan45.*

16 **El Niño-Southern Oscillation (ENSO) shapes extreme weather globally, causing myriad**
17 **socioeconomic impacts, but whether economies recover from ENSO events and how changes to**
18 **ENSO from anthropogenic forcing will affect the global economy are unknown. Here we show that**
19 **El Niño persistently reduces country-level economic growth, attributing \$4.1T and \$5.7T in global**
20 **income losses to the 1982-83 and 1997-98 events, respectively. Increased ENSO amplitude and**
21 **teleconnections from warming cause \$84T in 21st-century economic losses (discounted) in a middle-**
22 **of-the-road emissions scenario, but these effects are shaped by stochastic variation in the sequence**
23 **of El Niño and La Niña events. Our results highlight the sensitivity of the economy to climate**
24 **variability independent of warming and the possibility of future losses due to anthropogenic**
25 **intensification of such variability.**

26 As the leading mode of interannual climate variability, El Niño-Southern Oscillation (ENSO)
27 integrates a wide range of Earth system processes (1). El Niño events shift deep convection from the
28 western to the eastern Pacific, reorganizing global atmospheric circulation and shaping remote weather
29 through “teleconnections” (2, 3). The resulting temperature and hydroclimate extremes have many well-
30 documented impacts, including flooding (4, 5), crop losses (6, 7), and civil conflict (8). Many climate
31 models project that warming will increase El Niño amplitude (9, 10) and frequency (11), with potentially
32 devastating socioeconomic impacts (12).

33 Despite ENSO’s global impacts, however, empirical climate-economy studies have generally
34 focused on temperature and precipitation averages (13–18) or their variability (19), leaving the costs of
35 changes in modes of climate variability unquantified. While studies have shown that El Niño reduces
36 contemporaneous economic growth (20–22) and drives commodity price fluctuations (23–25), it remains
37 unclear if and for how long its impacts persist. Distinguishing between transient and persistent impacts on
38 economic growth is essential. Transient impacts (“level effects”) are quickly recovered, as an economy
39 rebounds to its original trajectory. Persistent impacts (“growth effects”) reduce an economy’s ability to
40 grow, compounding exponentially in time. Poor observational constraints on growth effects limit our
41 ability to understand the macroeconomic costs of ENSO and reduce this key uncertainty in climate
42 damage projections (26–28).

43 Here we estimate the effect of ENSO on economic growth historically and in the future,
44 accounting for the spatiotemporal heterogeneity of ENSO teleconnections. We define ENSO by the E-
45 index and C-index (29) (Fig. S1), metrics of El Niño and La Niña, respectively, that capture the nonlinear
46 feedbacks that drive ENSO (Methods). We define country-level teleconnections for each index (τ^E and τ^C)
47 using correlations between the indices and country-level temperature and rainfall (Methods, Fig. S2).
48 Teleconnections are strongest in tropical countries and weaker in the midlatitudes (Fig. 1a), consistent
49 with the physical responses of regional climate to tropical variability (30).

50 We use a distributed lag regression model to quantify the effect of ENSO on growth in national
51 Gross Domestic Product per capita (GDPpc) from 1960-2019. Departing from previous work (8, 21, 22),
52 we interact the E- and C-indices with teleconnections to allow the economic effect of ENSO to vary
53 smoothly as a function of teleconnection strength (31) (Methods). Our model compares economic growth
54 before and after El Niño events to assess their cumulative effects over time, allowing us to distinguish
55 growth from level effects (Methods). We focus on the five years following El Niño events, but also
56 evaluate effects for more than ten years and for La Niña events as well. We then couple these empirical
57 estimates with climate model projections to assess the future economic effects of changes to ENSO
58 amplitude and teleconnections.

59

60 **El Niño persistently reduces growth**

61 El Niño events persistently decrease economic growth (Fig. 1b). The magnitude of this effect is
62 determined by the strength of each country's E-index teleconnection. In Peru ($\tau^E = 1.18$), for example, a
63 1-standard-deviation (s.d.) El Niño event decreases growth by 1.3 percentage points (p.p.) in the year of
64 the event (95% confidence interval [CI]: 0.9 – 1.7 p.p.). Within five years, growth in Peru has declined by
65 6.2 p.p. (CI: 4.7 – 8.2) (Fig. 1b). By contrast, weakly teleconnected countries ($\tau^E < 0.5$) experience small
66 and uncertain effects (Fig. 1b). Interacting El Niño and teleconnections allows us to calculate marginal
67 effects for each country based on their τ^E value (Fig. 1c) and allows statistical significance to be
68 determined by uncertainty in the distributed lag model (hatching in Fig. 1c), rather than an ex ante
69 determination of “teleconnected” versus “non-teleconnected” countries. Some 56% of countries
70 experience significant declines in growth 5 years after an El Niño, averaging 2.3 p.p. Critically, the
71 increasing effect of El Niño with additional lags implies that most countries experience persistent growth
72 reductions after an event, not simply level effects from which they recover immediately (Fig. 1d, S3).

73 The negative growth effects of El Niño are robust to alternative methodological choices,
74 including different growth data, excluding strongly teleconnected countries, using alternative
75 teleconnection metrics, using more restrictive standard error clustering, and using the Niño3 index instead
76 of the E- and C-indices (SM Text, Figs. S3-S5). ENSO effects vary little over the 1960-2019 period,
77 indicating minimal treatment effect heterogeneity in time (Fig. S6). ENSO indices vary through time but
78 not space, raising the possibility that our model is confounded by time-varying global economic shocks.
79 While including year fixed effects could address this (32), our identification strategy does not easily
80 permit their addition (Methods). However, several alternative specifications illustrate that time-varying
81 confounders are not driving our results. The addition of country-specific trends to control for changes in
82 technology or demographics does not alter our results (Fig. S4) since ENSO is stochastic (33) and
83 measured by a detrended index. Alternative models that use a spatially varying country-level index of

84 ENSO or that discretize the sample into teleconnected and non-teleconnected groups allow us to include
85 both country and year fixed effects, and yield results as strong as our main estimates (Methods, Fig. S7).
86 Bootstrap resampling by year and estimates that drop each individual year or country, ensuring that single
87 years or countries are not driving the results, yield similar effects (Fig. S3). Finally, dropping the 1983
88 and 1998 events, which coincided with unrelated financial crises, reduces the magnitude of ENSO effects
89 by ~12%, but they remain statistically and economically significant (Fig. S4).

90 Our model reveals that El Niño effects can persist to 12 years or beyond, though a rebound begins
91 after ~10 years (Figs. S8, S9). Our focus on 5 lags in our results reflects a balance between tracing the
92 long-run response to ENSO and a concern for statistical power given the short observational record (Fig.
93 S8). Additional lags reduce the model degrees of freedom, leading to instabilities (Fig. S8). Data
94 simulations using a perfect model framework, where a permanent effect of El Niño is imputed to data,
95 demonstrate that models with many lags can yield insignificant coefficients due to the reduced sample
96 size and large number of parameters, even if the effect is known and permanent (Methods, Fig. S8).

97 Our empirical model includes both the E-index and C-index, allowing us to distinguish the effects
98 of eastern Pacific (EP) El Niño (where El Niños are strongest) and the central Pacific (CP) La Niña
99 (where La Niñas are strongest) (Methods). CP La Niña events have beneficial effects (Fig. S10), but they
100 are several times weaker than the negative effects of EP El Niño and are generally statistically
101 insignificant under more restrictive standard error clustering (Table S2). These results reflect the
102 skewness of ENSO itself, whereby EP El Niños tend to be stronger than both La Niñas and CP El Niños,
103 and are consistent with studies showing that La Niña's economic effect is small (21, 22).

104 Highly teleconnected countries most affected by ENSO are generally lower-income, tropical
105 countries (Fig. 1) (21). However, high-income countries still experience significant negative effects (Fig.
106 S4), consistent with work showing that these countries are impacted by extreme rainfall (18) and heat
107 (34), both of which ENSO affects. We also identify persistent losses across countries that experience
108 wetting and drying in response to El Niño (Fig. S4), as both anomalously low and high rainfall can be
109 damaging (18). More broadly, we emphasize that some regions can benefit from El Niño or be damaged
110 by La Niña. Our goal in this work is to estimate a globally generalizable response to ENSO. That our
111 findings are robust across multiple lines of country heterogeneity provides confidence that they are
112 generalizable, even if individual countries or regions within countries respond differently.

113

114 **Losses from historical El Niño events**

115 The persistent effect of ENSO implies that historical El Niño events have altered the income
116 growth of teleconnected countries, potentially generating large economic losses. Here we quantify the
117 costs of the two largest El Niño events in the last 60 years, the 1982-83 and 1997-98 events (Fig. 2).

118 Because an El Niño can trigger a subsequent La Niña (35), our analysis incorporates both the negative
119 effects of each El Niño and the benefits of the following La Niña (Methods). Furthermore, because these
120 events coincided with unrelated currency crises, we use a model excluding these two events to more
121 conservatively calculate their impacts (Fig. S4).

122 Consider strongly teleconnected Peru ($\tau^E = 1.18$): Its GDPpc declined in 1998 relative to 1997
123 and stagnated for three more years, before rising again (Fig. 2a). Given the 1997 financial crisis, Peru's
124 deviation from its growth trend in 1998 is not entirely attributable to ENSO, but Peru's economy would
125 have grown more quickly if the 1997-98 El Niño had not occurred (Methods). Income for the average
126 Peruvian would have been some \$1,246 greater five years later in 2003 absent the event (CI: \$853 –
127 \$1,793), an increase of 19% (Fig. 2a). Other tropical countries such as Ecuador, Brazil, and Indonesia
128 similarly lost anywhere from 5% to 19% of GDPpc due to the 1997-98 El Niño (Fig. S11).

129 We estimate global losses from the 1982-83 and 1997-98 events to be trillions of dollars each
130 (Fig. 2b, S11). Our estimates exceed previous ones because we account for ENSO's growth effects: one
131 study placed the total costs of the 1997-98 El Niño at \$36 billion (36). Our accounting has losses from the
132 1982-83 event rising to more than \$4.1T (CI: \$2.3T – \$6T) by 1988. Similarly, the costs of the 1997-98
133 event reached \$5.7T (CI: \$2.3T – \$9.2T) by 2003. The greater costs of the 1997-98 event result both
134 because it was a stronger El Niño and because the global economy was larger. Absent the compensating
135 benefits of the subsequent La Niñas, the 1983 (1998) event would have led to global losses of \$4.4T
136 (\$8T) (Fig. 2b). By considering overall GDP, incorporating growth reductions following the event, and
137 including all countries in a single framework, our findings show that estimates focusing on physical asset
138 losses in low-income countries have strongly underestimated the global economic toll of El Niño.

139

140 **Climate model projections of ENSO**

141 ENSO's persistent effect raises the question of how it will shape the global economy with global
142 warming. Using climate model simulations from the sixth phase of the Coupled Model Intercomparison
143 Project (CMIP6) that skillfully represent the skewness of eastern Pacific SSTs, we analyze projected
144 changes to ENSO under four Shared Socioeconomic Pathways (SSPs) (Methods).

145 El Niño amplitude and teleconnections are projected to increase relative to the historical period in
146 CMIP6 (Fig. 3). Such a response is not scenario-dependent, likely due to the influence of internal climate
147 variability on forced ENSO changes (37–39). Amplitude increases by 5 – 21% in the median across
148 scenarios (Fig. 3a), a function of stronger wind-ocean coupling in the eastern Pacific (9, 12). Similarly,
149 global mean teleconnections increase by 4 – 15% (Fig. 3b), consistent with a more energetic atmospheric
150 response to El Niño (40, 41). Despite the shared response in amplitude and teleconnection changes,

151 internal variability (proxied by multiple realizations from each model) can vary these responses by
152 upwards of 60% (Fig. 3a, b, lower lines).

153 Beyond amplitude and teleconnection changes, climate projections also differ in their E-index
154 time series, with implications for future damages. Due to ENSO's sensitivity to initial conditions (37–39)
155 and multidecadal variability (42, 43), a wide range of E-index values across models and scenarios can
156 occur in a given year, even controlling for amplitude (Fig. 3c, Fig. S12). For example, Figure 3c shows
157 two SSP2-4.5 simulations with similar amplitude changes and E-index skewness but different sequences
158 of eastern Pacific El Niños and La Niñas. As quantified by the sum of the E-index over the 21st century,
159 MIROC-ES2L r6i1p1f2 experiences strong El Niño events, especially early in the 21st century, while
160 CESM2-WACCM r3i1p1f1 is dominated by La Niña events. Such differences in the sequence of ENSO
161 shape projected damages, as an El Niño-dominated time series yields greater damages than a La Niña-
162 dominated one due to their differential effects (e.g., Fig. S10). Crucially, because El Niños are stronger
163 than La Niñas, the long-run expectation from increased ENSO amplitude is net economic losses.

164 We combine these projections with our empirical estimates to quantify the economic effects of
165 changes in ENSO. We use the SSPs as baselines against which we calculate country-level growth changes
166 based on projections of ENSO amplitude and teleconnection change (Methods). Because the effect of
167 ENSO may rebound after ~14 years and we cannot confidently identify truly permanent growth effects
168 due to the short observational record (Fig. S8), we make the conservative choice to allow economies to
169 recover from ENSO events after 14 years in our projections, though we also show damages under the
170 assumption that they are permanent (Methods, Fig. S13).

171

172 **Economic impacts of future ENSO changes**

173 Anthropogenic changes to El Niño amplitude and teleconnections will likely cause substantial
174 economic losses over the 21st century (Fig. 4). Under a 2% discount rate (44) and a realistic emissions
175 trajectory (45) (SSP2-4.5), the median cumulative 2020-2099 global losses are \$84T (Fig. 4a), a ~1%
176 reduction in global economic output over the 21st century. In all four scenarios, median losses exceed
177 \$18T and damages are negatively skewed, consistent with the asymmetry in ENSO itself.

178 The range of these projected losses is large. Under SSP2-4.5, the 95% range spans losses of
179 \$453T to benefits of \$80T (we write this CI as -\$453T – +\$80T) across 86,000 combinations of 86
180 simulations and 1,000 regression bootstraps (Fig. 4a). Reducing the discount rate to 1% amplifies median
181 losses under SSP2-4.5 to \$130T (-\$687T – +\$130T), while increasing it to 5% diminishes losses to \$26T
182 (-\$162T – +\$34T). The extreme end of these ranges implies a ~5% reduction in global economic output
183 over the remaining century. In highly teleconnected countries, changes to ENSO are associated with
184 GDPpc reductions of >1% per year, though uncertainty is high even in these countries (Fig. S11).

185 Despite this range across models, scenarios, and internal variability, increases in ENSO amplitude
186 and teleconnections are systematically related to greater economic losses (Fig. 4b, c). Each 1% increase in
187 ENSO amplitude is associated with \$4.1T in additional discounted losses over the 21st century ($p <$
188 0.001), and each 1% increase in teleconnections is associated with \$6.3T in additional losses ($p <$ 0.001).
189 These findings build upon previous projections of changes in ENSO amplitude (9, 11) and
190 teleconnections (40, 41), demonstrating tangible, global socioeconomic effects of these physical changes.

191 These relationships, however, are heterogeneous, as the largely stochastic sequence of El Niños
192 and La Niñas going forward shapes the direction and magnitude of damages. Simulations with E-index
193 sums greater than 0 (El Niño-dominated time series) exhibit a strong negative relationship between ENSO
194 amplitude increases and damages (Fig. 4b, red dots), but the opposite is true for La Niña-dominated time
195 series (Fig. 4b, blue dots). The same pattern holds for teleconnection changes (Fig. 4c). Critically,
196 because ENSO is asymmetric and El Niños are generally stronger than La Niñas, there are many more El
197 Niño- than La Niña-dominated time series. On average, therefore, increases in ENSO amplitude and
198 teleconnections produce large global economic losses.

199 Alternative methodological choices, including incorporating changes in the C-index or holding
200 teleconnections constant, alter the magnitude of losses but do not change the core result of negative
201 ENSO-driven damages with warming (Fig. S13). Using only one realization from each model increases
202 uncertainty across scenarios (Fig. S13b), suggesting the importance of large ensembles to effectively
203 capture ENSO variability (38). Assuming that damage persistence is permanent substantially increases the
204 magnitude and uncertainty in projected damages (Fig. S13d). Finally, controlling for country-level
205 average temperature and rainfall in our regression does not alter the effect of ENSO (Fig. S14), meaning
206 the effects we identify are distinct from damage projections using average temperature (13). ENSO may
207 affect sub-national or sub-annual extreme temperature or rainfall, as well as other hazards such as
208 drought, all of which may have independent economic impacts (18, 46, 47).

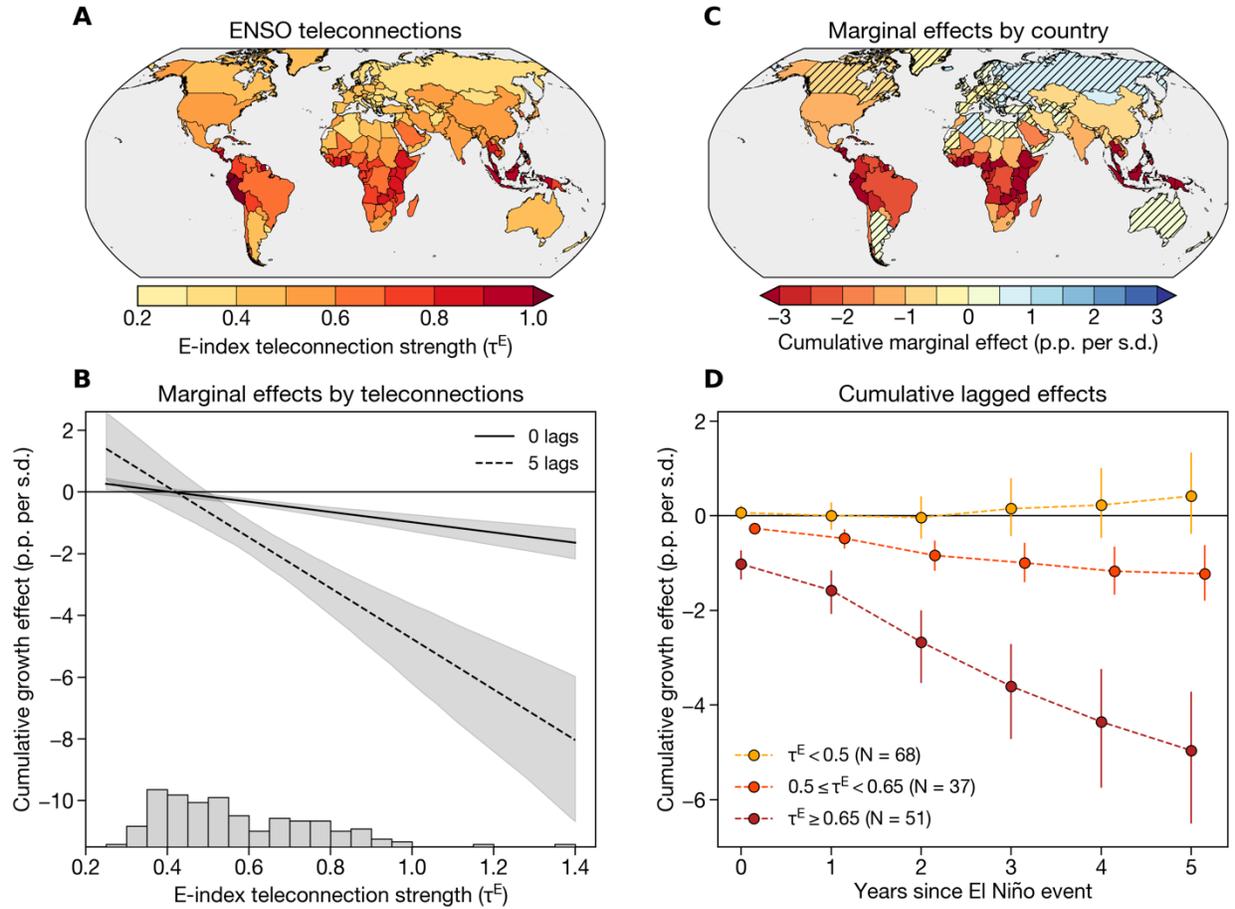
209 Our findings have implications for climate mitigation and adaptation. All else being equal,
210 increased ENSO amplitude and teleconnections will generate major economic losses not currently
211 included in assessments of climate damages or mitigation benefits. However, the facts that (1) ENSO-
212 driven damages do not depend strongly on future climate scenario (Fig. 4a) and (2) a range of outcomes
213 are possible due to uncertainty in the unique ENSO sequence the world experiences going forward (Fig.
214 4b, c) together imply that emissions reductions alone are insufficient to protect economies from El Niño.
215 While mitigation remains the most effective means to blunt the catastrophic impacts of anthropogenic
216 warming (48), our findings simultaneously raise the priority of climate adaptation and resilience efforts.
217 Improved disaster risk management could reduce ENSO-driven damages (49), and scientific investments

218 in ENSO early warning and decadal climate prediction could reduce the uncertainty in projections of
219 these damages.

220

221 **Conclusion**

222 Our finding that El Niño has a persistent effect on economic growth has four key implications:
223 Firstly, it demonstrates that, independent of warming, economic growth is highly sensitive to climate
224 variability. The impacts we identify demonstrate that the local extreme conditions associated with ENSO
225 integrate into a globally persistent macroeconomic effect, implying large and previously underestimated
226 costs of historical El Niño events. Secondly, our results demonstrate that future changes to ENSO may
227 increase the macroeconomic costs of warming. Previous climate-economy studies have not incorporated
228 changes in climate variability, and we show that this omission has hidden a potentially major cost of
229 rising temperatures. Thirdly, stochastic variation in ENSO could result in either losses or benefits from
230 warming even in the face of enhanced ENSO amplitude and teleconnections, emphasizing the importance
231 of investing in ENSO prediction, particularly on decadal time scales (42). Lastly, these findings together
232 suggest that while climate mitigation is essential to reduce accumulating damages from warming, it is
233 imperative to devote more resources to adapting to El Niño, as economies are more vulnerable to climate
234 variability than previously understood.



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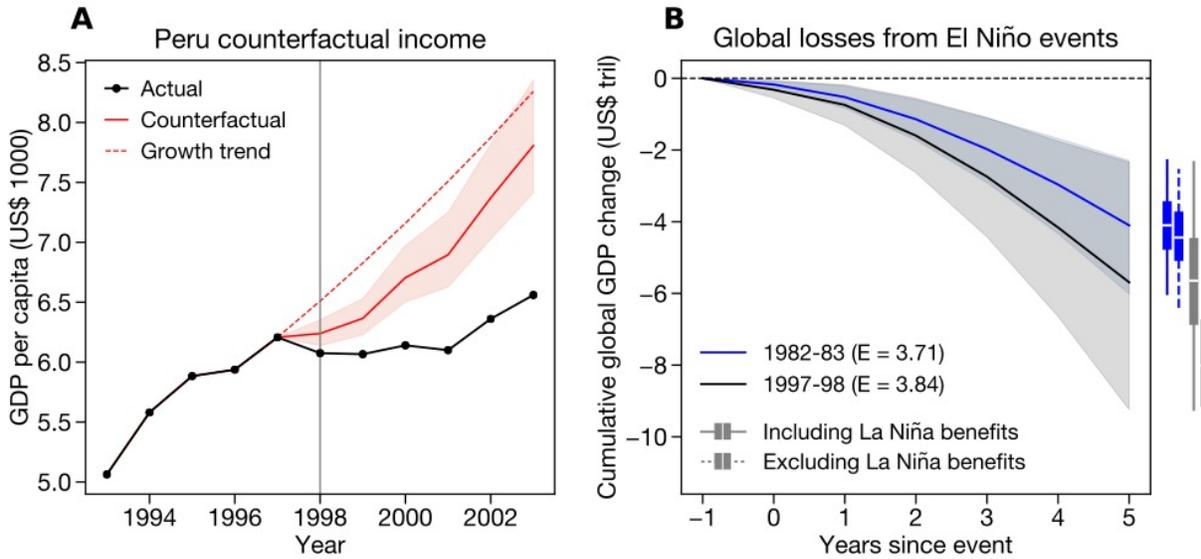
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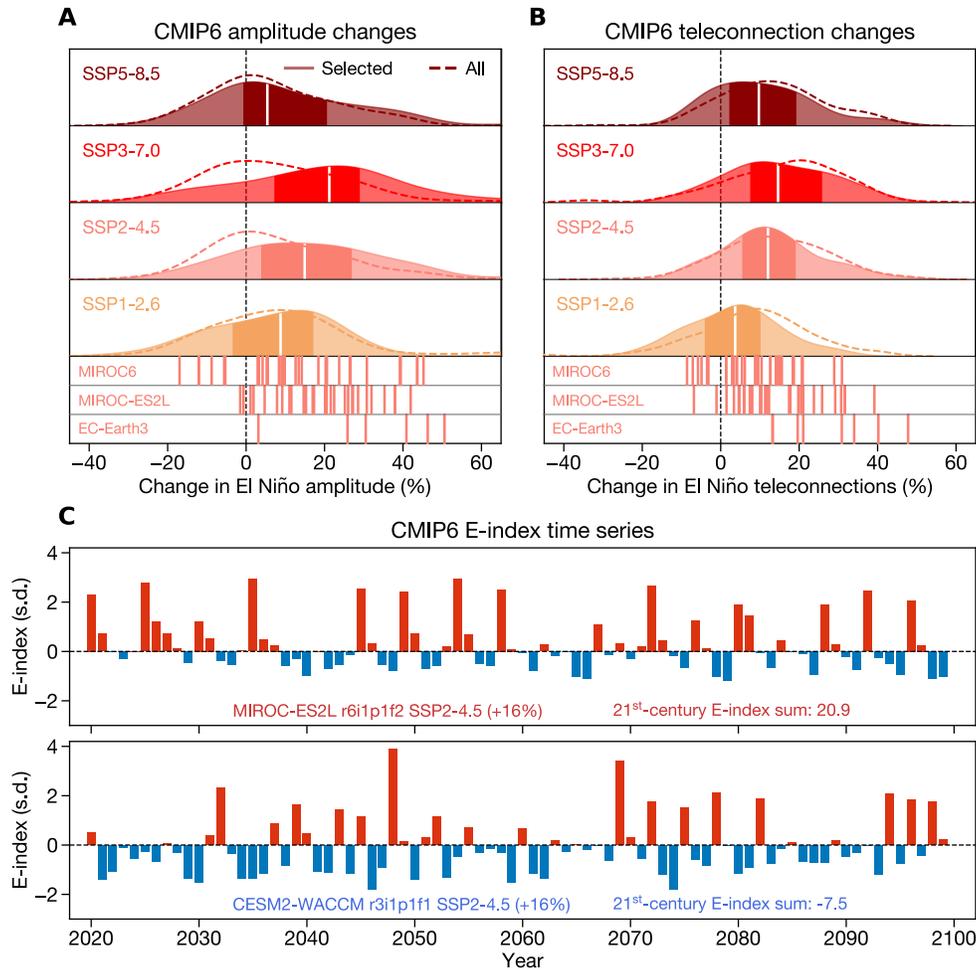
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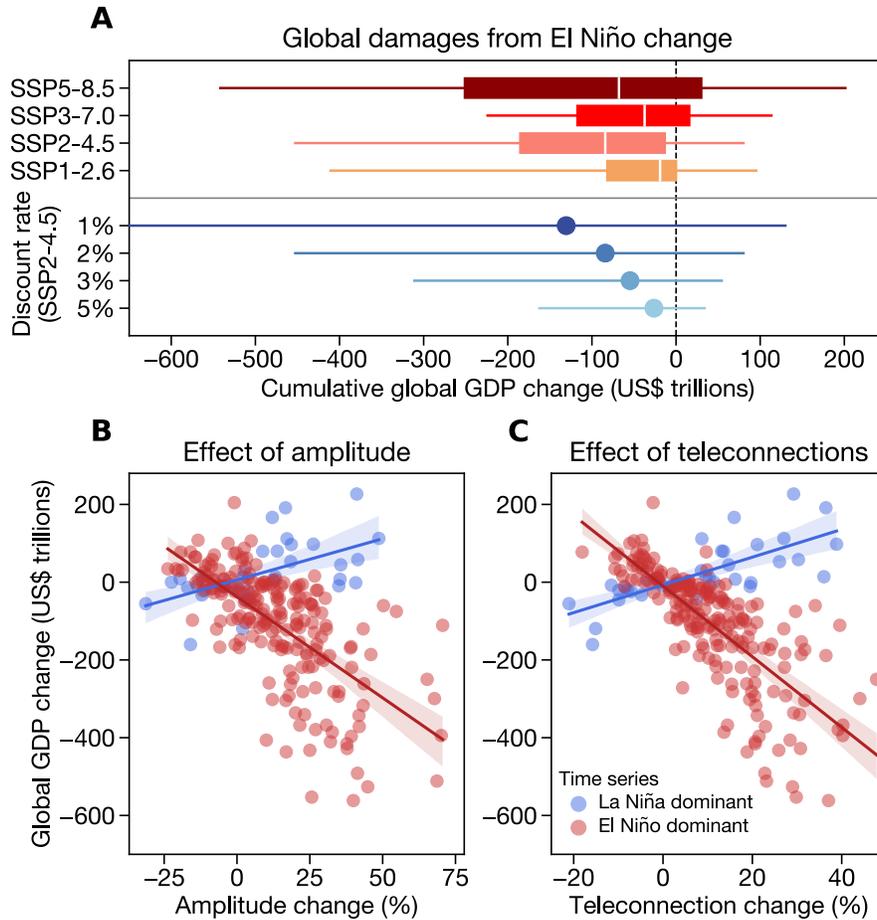
Fig. 1 | Teleconnections mediate the effect of El Niño on economic growth. **A)** Country-level ENSO teleconnections, calculated as the sum of the absolute value of the correlation coefficients between the E-index and monthly country-level temperature and precipitation (Methods). **B)** Marginal effects of El Niño on economic growth across teleconnection values in year of the event (0 lags, solid line) and the fifth year after the event (5 lags, dashed line). Black line shows the mean and shading shows 95% confidence intervals from bootstrap resampling (Methods). Lower histogram shows the density of teleconnection values in the sample. **C)** Cumulative 5-lag effect of El Niño on economic growth for each country. Hatching denotes countries whose effects are not distinguishable from zero (i.e., they fall on a location on the x-axis in (B) where the shading includes zero). **D)** Cumulative effects of El Niño over time, beginning with the year of the event (year 0) and accumulating to the fifth year after the event (year 5). Countries are grouped into three bins according to their teleconnection strength, with “N” denoting the number of countries in each bin. Dots show averages and bars show 95% confidence intervals.



248
 249 **Fig. 2 | Damages from extreme El Niño events.** **A)** GDP per capita (GDPpc) in Peru before and after the
 250 1997-98 El Niño event. Black line shows actual GDPpc, red line shows the average counterfactual
 251 GDPpc across regression bootstrap samples (Methods), and red shading shows 95% confidence interval.
 252 Dashed line shows GDPpc if Peru had maintained its average growth rate from the 5 years preceding the
 253 event. **B)** Cumulative global GDP change for the 5 years after the 1982-83 (blue) and 1997-98 (black) El
 254 Niño events. Center line shows the mean and shading shows the 95% confidence intervals across
 255 regression bootstrap samples. Global GDP change is only calculated for countries with statistically
 256 significant marginal effects (Fig. 1c). Text in legends denotes the DJF-average E-index in the
 257 corresponding years. Boxplots at right show cumulative global GDP change when including the benefits
 258 of the following La Niña events (solid lines) and excluding those benefits (dashed lines). All dollar values
 259 are in constant 2017 prices.



260
 261 **Fig. 3 | Climate model projections of ENSO.** Change in ENSO amplitude (**A**) and global mean
 262 teleconnection strength (**B**) between 1940-2019 and 2020-2099 for an ensemble of CMIP6 simulations
 263 from four SSP experiments. In both panels, dashed density lines show changes from all simulations and
 264 solid density plots show amplitude changes from selected high-skill simulations used in the analysis
 265 (Methods). Vertical lines below density plots denote amplitude changes from the individual realizations
 266 of three models (MIROC6, MIROC-ES2L, and EC-Earth3), all drawn from the SSP2-4.5 experiment,
 267 illustrating the wide range of amplitude and teleconnection changes possible from internal variability
 268 alone. **C**) E-index time series from two example simulations with similar amplitude increases: MIROC-
 269 ES2L r6i1p1f2 (top) and CESM2-WACCM r3i1p1f1 (bottom), both from the SSP2-4.5 experiment. Red
 270 bars denote eastern Pacific El Niño ($E > 0$) and blue bars denote eastern Pacific La Niña ($E < 0$). Left
 271 inset text in each panel denotes the model information and amplitude change. Right inset text denotes the
 272 sum of each E-index time series over the 21st century (2020-99), with positive values indicating that the
 273 time series contains more El Niños than La Niñas and negative values indicating the opposite.



274

275 **Fig. 4 | Global economic impacts of changes in El Niño amplitude and teleconnections. A)** Boxplots
 276 show the cumulative global GDP change in each scenario under a 2% constant discount rate. Colors
 277 correspond to the scenario colors in Fig. 3. In each boxplot, white line denotes the median, box spans the
 278 first and third quartiles, and whiskers span the 95% range. Lower blue lines denote global economic
 279 losses under SSP2-4.5 and a range of discount rates. Dot denotes the median and lines span the 95%
 280 range. **B, C)** Cumulative global GDP change due to changes in ENSO amplitude (**B**) and teleconnections
 281 (**C**) with a 2% discount rate, with each dot corresponding to one climate model simulation. Simulations
 282 are pooled across all four scenarios. Red dots denote simulations in which the 21st-century E-index sum is
 283 greater than 0 (El Niño-dominated time series), while blue dots denote simulations in which the sum is
 284 less than 0 (La Niña-dominated time series). Red and blue regression lines and 95% CIs are drawn
 285 separately for each subset of simulations.

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Supplementary Materials for
Persistent Effect of El Niño on Global Economic Growth

Christopher W. Callahan and Justin S. Mankin

correspondence to: Christopher.W.Callahan.GR@dartmouth.edu

This PDF file includes:

- Materials and Methods
- Supplementary Text
- References (50-84)
- Figs. S1 to S17
- Tables S1 to S7

310 **Materials and Methods**

311 Data

312 We use observational climate data from multiple sources: Monthly mean sea surface
313 temperatures (SST) from the HadISST dataset (50), monthly mean atmospheric temperatures
314 from the Berkeley Earth dataset (51), and monthly total precipitation data from the Global
315 Precipitation Climatology Center (52). Temperature and precipitation are aggregated to
316 population-weighted country-level means using year-2000 population data from the Gridded
317 Population of the World (53). We use population weighting to ensure that the spatial aggregation
318 captures climate fluctuations that affect people and economic activity.

319 We use country-level economic data from the Penn World Tables version 10.0 (54),
320 specifically Gross Domestic Product (“RGDPNA”) (in 2017-equivalent dollars) and population
321 (“POP”) for all countries of the world. GDP per capita (GDPpc) is calculated as GDP divided by
322 population. Growth for each year is calculated as the fractional GDPpc change relative to the
323 previous year. Because macroeconomic data may contain measurement error (55), we also repeat
324 the analysis using data from the World Bank World Development Indicators (56), finding similar
325 results (Fig. S4).

326 The time period of analysis for both the teleconnection calculations and regression
327 analysis is 1960-2019, so all observational economic and climate data is limited to that time
328 period.

329 Climate model data come from the sixth phase of the Climate Model Intercomparison
330 Project (57) (CMIP6). We use monthly SST, monthly atmospheric temperature, and daily
331 precipitation data over 1850-2099 from the historical experiment and the four Tier 1 experiments
332 from the Scenario Model Intercomparison Project (58). These four experiments—SSP1-2.6,
333 SSP2-4.5, SSP3-7.0, and SSP5-8.5—span a range of plausible policy futures, from aggressive
334 mitigation (SSP1-2.6) to high emissions (SSP5-8.5) (58, 59). Global mean temperatures rise by
335 ~1.2 °C by 2081-2100 relative to 1995-2014 in the SSP1-2.6 scenario, 2.1 °C in SSP2-4.5, 3.2
336 °C in SSP3-7.0, and 4 °C in SSP5-8.5 (59). Not all models have data available for each
337 experiment, so differences across the experiments are due both to differences in forcing and
338 differences in the sampling of model structure (Tables S3-S6). All climate model data is
339 regridded to a 2°-by-2° grid, using bilinear interpolation from Python’s “xarray” package (60).

340

341 ENSO indices

342 We use the “E-index” and “C-index” to represent ENSO (9, 29, 37, 61, 62). The E-index
343 represents eastern Pacific El Niño events and captures the nonlinear processes that generate
344 skewness in eastern Pacific SSTs, whereby El Niño events are stronger than La Niña events (9,
345 29). The E-index is a combination of the first two principal components (PCs) of an empirical
346 orthogonal function (EOF) analysis applied to Pacific SSTs (37) over 20 °S – 20 °N and 140 °E
347 – 80 °W, specifically as $E = (PC1 - PC2)/\sqrt{2}$. We calculate the E-index in observations using
348 linearly detrended SST anomalies referenced to 1960-2019 long-term monthly means. We then
349 average the E-index over winter (December-February, DJF), to focus on the season in which
350 ENSO peaks (63); the E-index in year t is therefore defined as the average of the December E-
351 index from year $t-1$ and the January and February indices from year t .

352 The C-index (29) is a companion index to the E-index and is calculated as $C = (PC1 +$
353 $PC2)/\sqrt{2}$. The C-index represents central Pacific La Niña and El Niño events, where La Niña
354 events tend to be stronger than El Niño events. Positive E-index values represent an eastern
355 Pacific El Niño event and negative C-index values represent a central Pacific La Niña event. The

356 E-index and C-index are orthogonal by construction (29), allowing us to include them both in a
357 regression model without a concern for collinearity.

358 To assess the sensitivity of our results to these indices, we also calculate the Niño3 index,
359 defined as linearly detrended SST anomalies averaged over 5 °S – 5 °N and 150 °W – 90 °W.
360 The Niño3 index yields similar, though slightly weaker, results to the E-index (Fig. S4) since it
361 corresponds to eastern Pacific conditions but does not distinguish the spatial structures of El
362 Niño and La Niña.

363 We calculate the DJF E- and C-index similarly in the CMIP6 models, using quadratically
364 detrended (9) SST anomalies referenced to monthly means from 1850-2014.

365 Country-level ENSO teleconnections

366 Our analysis incorporates a country-specific teleconnection metric to quantify heterogeneity
367 in growth responses according to a country’s geophysical connection to ENSO. To calculate the
368 teleconnection, we first standardize monthly country-level mean temperature and total
369 precipitation by subtracting the long-term (1960-2019) monthly mean and dividing by the long-
370 term monthly standard deviation. We then linearly detrend these standardized anomalies
371 separately for each month to remove the effects of warming and low-frequency climate
372 variability.
373

374 Next, we correlate these standardized temperature and precipitation time series with the DJF
375 E-index separately for each month m and each country i . El Niño events begin and grow in year
376 $t-1$, peak in the winter, and then decay in the spring and summer of year t , so we allow the DJF
377 E-index to affect both the preceding (beginning just after the “spring predictability barrier” in
378 June of $t-1$) and following years (ending in August of year t) (Fig. S1). We use partial
379 correlations to control for precipitation when analyzing temperature and vice versa to control for
380 the covariance between temperature and precipitation.

381 This calculation yields a distribution of 15 correlation coefficients (one per month from
382 June of year $t-1$ through August of year t) for each country, separately for temperature and
383 precipitation. We then take the three-month running mean of these coefficients across the 15
384 months to smooth out random variation and account for multiple months of exposure to ENSO.
385 Finally, we take the maximum (absolute) correlation coefficients from these running means for
386 both temperature and precipitation and add them together to calculate each country’s E-index
387 teleconnection τ^E . We use absolute values to allow the distinct effects of temperature and
388 precipitation teleconnections to be additive, but our results are robust to considering both
389 positive and negative precipitation teleconnections separately (Fig. S5).

390 This teleconnection metric estimates the degree to which each country’s climate is
391 influenced by ENSO, accounting for: (1) the effects of both temperature and precipitation; (2)
392 multiple sustained months of exposure to ENSO; and (3) the varied timescales on which
393 teleconnections may manifest. Additionally, this strategy allows teleconnections to be defined
394 continuously rather than separating teleconnected and non-teleconnected countries based on
395 arbitrary significance thresholds (8) or previously defined climate zones (21, 22). Fig. S2 shows
396 the steps in this teleconnection calculation, and we perform the same analysis with the C-index to
397 calculate C-index teleconnections (τ^C).

398 Econometric analysis

400 The goal of our analysis is to quantify the multi-year effect of ENSO on economic growth.
401 This task requires us to separate ENSO from the other constant and time-varying factors that

402 affect economic growth. We use a distributed lag regression model, estimated with Ordinary
 403 Least Squares, to estimate the effects of eastern Pacific El Niño (the E-index) and central Pacific
 404 La Niña (the C-index) on growth:
 405

$$g_{it} = \sum_{L=0}^j [\beta_L E_{t-L} + \Theta_L E_{t-L} * \tau_i^E + \Phi_L C_{t-L} + \Psi_L C_{t-L} * \tau_i^C] + \mu_i + \epsilon_{it} \quad (1)$$

407
 408 Here, g refers to growth in country i in time t , E refers to the E-index in year t , and C refers
 409 to the C-index in year t . μ is a country fixed effect, which controls for average differences
 410 between countries such as geography and ensures that our results are identified using within-
 411 country variation in growth. L is the lag at which the coefficient is estimated. The interactions of
 412 E with τ^E and C with τ^C allow the effect of ENSO to differ between countries based on how
 413 strongly coupled each country's climate is to ENSO.

414 The inclusion of lagged terms from years L to j allows us to distinguish between level and
 415 growth effects on the economy. If the effect of El Niño only falls on income levels, then a shock
 416 in year t will be recovered in year $t+1$ as countries rebound to their original income trajectory,
 417 meaning that year $t+1$ will see an abnormally high growth rate. If, instead, El Niño affects the
 418 underlying capacity of the economy to grow, then the years following an event should show
 419 either persistent declines in growth or no change. As such, our analysis focuses on the
 420 cumulative coefficients Ω , which represent the accumulated effect of ENSO in the years after an
 421 event. The interaction of E with country-specific teleconnections τ^E allows us to calculate unique
 422 cumulative effects for each country i and lag length L :
 423

$$\Omega_{iL} = \sum_{L=0}^j [\beta_L + \Theta_L * \tau_i^E] \quad (2)$$

425
 426 If Ω_{iL} is indistinguishable from zero, then we cannot reject the hypothesis that El Niño has
 427 only level effects; growth effects are identified if Ω_{iL} is significantly different from zero ($p <$
 428 0.05). Note that the E-index is not highly correlated with itself across lag lengths (Table S7),
 429 meaning that including multiple lags in a single model should not generate multicollinearity.

430 The identifying variation in our model comes from stochastic and unpredictable (33, 64)
 431 shifts in SSTs from year to year, along with the differential effects of those SSTs depending on
 432 teleconnection strength. The E- and C-index are constant throughout space within a given year,
 433 raising the concern that other time-varying confounders could be correlated with ENSO and
 434 generating spurious results. A typical strategy in empirical climate-economy studies is to include
 435 both unit and time fixed effects in regression models (32), which separates local weather
 436 variation from both time-invariant average conditions and global time-varying shocks. However,
 437 because the E- and C-index terms in Eqn. 1 would be collinear with the year fixed effect, we
 438 cannot estimate our main specification with year effects.

439 We do, however, show results from several alternative models that separate the influence of
 440 ENSO from time-varying confounders. First, adding linear or linear and quadratic country-level
 441 time trends to control for secular trends in technology or demographics does not alter our results
 442 (Fig. S4). Second, bootstrap resampling by year permutes the years in the regression model and
 443 ensures that no single year has a disproportionate influence on the results (Fig. S3). Third,

444 dropping 1983 and 1998 from our data, which were major El Niño events that coincided with
445 financial crises in tropical countries, reduces the magnitude of the effects we find by ~12% but
446 does not alter their statistical or economic significance (Fig. S4). Fourth, we define a unique
447 spatiotemporally varying ENSO index for each country and year by multiplying E_t by τ_i^E .
448 Because this index differs across countries within years, we can estimate the model with country
449 and year fixed effects, and we find negative effects that exceed the results of our main model
450 (Fig. S7). For example, this model predicts that Peru experiences an 8.7-p.p. decline in growth
451 five years after an El Niño, compared to 6.2 p.p. from our original model. Finally, we estimate a
452 discretized version our main model, where we defined “untreated” countries as countries with τ_i^E
453 < 0.5 and “treated” countries as countries with $\tau_i^E > 0.5$. This allows us to estimate the model
454 with country and year fixed effects, interpreting the discretized interaction term as the effect of
455 ENSO on treated countries. In this case, we find that treated countries experience >3-p.p.
456 declines in growth five years after El Niños, which exceeds the 2.3-p.p. average loss for
457 countries with $\tau_i^E > 0.5$ from our main model (Fig. S7). The inclusion of year fixed effects in
458 these latter two models, along with the other checks we show, supports our conclusion that our
459 results are not driven by time-varying confounders.

460 We estimate confidence intervals by bootstrapping ($N = 1,000$), with countries resampled
461 from a uniform distribution with replacement. Countries are sampled as a block to account for
462 within-country autocorrelation (65). However, alternative bootstrapping schemes yield similar
463 results, such as sampling by year globally or within continents to account for spatial correlation
464 in growth, sampling by continent to account for simultaneous spatial and temporal correlation,
465 and sampling by five-year blocks to account for spatial and short-term temporal correlation (65)
466 (Fig. S3). Multiple forms of clustered parametric standard errors, which are robust to both
467 spatiotemporal autocorrelation in errors and heteroskedasticity across clusters, do not reduce the
468 statistical significance of our results (Table S1).

469 We remove growth values from our sample that are above 18% or below -18%,
470 approximately the 3σ range. We drop 138 values because of this choice, less than 2% of the
471 sample. Including these values does not reduce the average effect, but it does increase the
472 uncertainty (Fig. S4), so we drop these outliers while noting that our results would be similar if
473 we included them.

474 When we estimate separate responses for high-income and low-income countries (Fig.
475 S4), we use the World Bank’s income classifications, grouping low and lower-middle income
476 countries together as well as high and higher-middle income countries. Again, the results accord
477 with our main model.

478 Other time series analysis tools have been used to assess the effect of ENSO such as
479 vector autoregression (VAR) models (20, 23–25) or local projections (20). We use a distributed
480 lag (DL) model for two reasons. Firstly, DL models have been widely used in the empirical
481 climate-economy literature (13, 15, 66, 67), so our approach is consistent with this work.
482 Secondly, VAR models are primarily used in macroeconomic settings where endogeneity is at
483 issue (68). Because ENSO is plausibly exogenous to country-level growth rates, we adopt the
484 more parsimonious DL model.

485

486 Synthetic data simulations

487 Estimating the effect of El Niño with models that include 14 or more lags results in
488 unTable coefficients and confidence intervals that include zero (Fig. S8). Two plausible
489 interpretations of this result are: (1) that there is no statistically significant growth effect of El

490 Niño after 14 years; or (2) that there is a permanent growth effect, but models with many lags
491 cannot confidently identify this effect due to the reduced sample size and increased number of
492 parameters being estimated simultaneously.

493 To examine this issue, we use a perfect model framework where we impute a known El
494 Niño effect to synthetic growth data and then estimate the regression on that data to assess
495 whether we can recover the effect. We construct growth as the combination of a first-order
496 autocorrelated process (AR(1)) with Gaussian noise of mean 0 and s.d. 0.05, a linear trend
497 randomly chosen from a Gaussian distribution of mean 0 and s.d. 0.2 (in p.p. per year), and an El
498 Niño effect. The AR(1) coefficient is set to 0.1, within the range of AR(1) coefficients from the
499 data, and the distribution of trends we choose from is also similar to the distribution of country-
500 level growth trends from the data (Fig. S15).

501 We then create a “true” effect of ENSO on growth and attempt to recover it with the DL
502 model. This predetermined ENSO effect is ultimately arbitrary, but we choose country-level
503 effects that are similar in magnitude to the effects we find in our main regression. We set these
504 effects to accumulate over the first 5 years and plateau at that 5-year value permanently. The
505 non-interacted effect of E is set to sum to 3 p.p. per s.d. and the interaction coefficient with τ is
506 set to sum to -6 p.p. per s.d., meaning that a country with $\tau^E = 1.0$ experiences a cumulative
507 effect of -3.0 p.p. per s.d. ($3 + 1.0 \cdot -6$).

508 We then fit Eqn. 1 using this synthetic growth data and the actual E-index and τ^E values,
509 using between 5 and 18 lags in the regression (beyond 18 lags, the coefficients become
510 undefined as the degrees of freedom decrease). We repeat this entire process 1,000 times for
511 each number of lags, keeping the set El Niño effect constant. Fig. S8 shows the results from
512 these estimations for one example teleconnection value ($\tau^E = 1.0$). These models are generally
513 unbiased, with the central estimate matching the imputed effect. However, confidence intervals
514 steadily grow as lags are added. With 14 or more lags, the coefficients become statistically
515 insignificant. These results demonstrate that even with a known permanent effect of El Niño,
516 estimating additional lag terms induces sufficient uncertainty to yield insignificant coefficients.
517 To assume that El Niño has no effect in the 14-lag model therefore risks a Type II error. That
518 being said, as a conservative choice in our historical attribution and in our damage projections,
519 we only allow the effects to be partially persistent rather than permanently persistent (see
520 *Economic damages from changes to ENSO*). In our attribution of the costs of the 1982-83 and
521 1997-98 events, we estimate costs accumulating to 5 years after the event. In our projections, we
522 allow effects to accumulate to 14 years, the maximum length we can confidently identify effects
523 from the observational data (Fig. S8). In a sensitivity test, we allow the effects to be permanent
524 (Fig. S13).

525 Economic damages from historical extreme El Niño events

527 The regression coefficients derived from Eqn. 1, β and θ , provide estimates of the change
528 in economic growth for a 1-s.d. change in the E-index. These coefficients can then be applied to
529 actual and hypothetical E-index time series to calculate the growth effects of specific historical
530 El Niño events. Here we focus on the two major El Niño events of 1982-83 and 1997-98. We
531 develop “counterfactual” E-index time series wherein these events did not occur by setting the
532 corresponding E-index values (1983 and 1998) to zero. We then apply the regression coefficients
533 to the actual and counterfactual time series to calculate the growth difference between them over
534 the five years after the event. Formally, if E^O represents the observed E-index in the year of the

535 event (t), and E^{CF} represents the counterfactual E-index in that year, we calculate the growth
536 change in country i from year t through year $t+L$ as:

537
538
$$\Delta g_{i(t+L)} = [\beta_L E_t^{CF} + \Theta_L E_t^{CF} * \tau_i^E] - [\beta_L E_t^O + \Theta_L E_t^O * \tau_i^E]$$
 (3)
539

540 We add these growth change values to the observed growth data, yielding a counterfactual
541 growth time series, and we integrate counterfactual growth to calculate counterfactual income
542 from the year of the event to 5 years after the event. Losses due to each event are calculated as
543 the difference between observed and counterfactual income. Details of this procedure can be
544 found in Diffenbaugh and Burke (69).

545 Note that E^{CF} is zero in our analysis, so the first bracketed term on the right-hand-side of
546 Eqn. 3 is zero, but we provide the full equation because it generalizes to other counterfactual E-
547 index values.

548 The above analysis only incorporates reductions in growth due to the El Niño events.
549 However, because El Niño events can dynamically trigger La Niña events (35), which have
550 beneficial effects (Fig. S10), a full accounting of the effects of El Niño should incorporate these
551 offsetting beneficial events. The 1982-83 El Niño may have triggered the La Niña of 1984-85
552 (while the C-index was only -0.07 in 1984, it was -1.1 in 1985), and the 1997-98 El Niño may
553 have triggered the major La Niña of 1999-2000 (the C-index was -2.1 in 1999 and -2.0 in 2000).
554 We incorporate these beneficial effects for both El Niño events by setting the C-index values for
555 the following two years (i.e., 1999 and 2000 in the case of the 1998 El Niño) to zero and
556 calculating the growth difference between the actual and counterfactual C-index time series. The
557 total growth change over the five years following the El Niño event is therefore the reduction due
558 to the El Niño event plus the increase due to the following La Niña events.

559 For both events, we limit our analysis to countries with continuous GDPpc data since
560 1982 to ensure that the same countries are included in both calculations. This restriction means
561 that nations with short GDPpc records (e.g., post-Soviet nations like Ukraine) are not included in
562 these calculations.

563 Climate model selection

564 Many climate models do not realistically represent the physical processes that drive
565 ENSO (70–72). To ensure that our projections are physically realistic, we filter the simulations
566 we use based on criteria set out in previous studies (9, 37, 72). We calculate a parameter known
567 as α from each model, which is the quadratic coefficient on the relationship between the first and
568 second principal components from the EOF analysis used to calculate the E-index and C-index
569 (72) (see *ENSO indices*).

570 The observed value of α is -0.34, indicating a strong nonlinearity in the principal
571 component space and a strong differentiation between eastern Pacific and central Pacific El Niño
572 events. Models which simulate an α value closer to the observed value also more effectively
573 represent the variance and skewness in SST anomalies, as well as the distinct eastern and central
574 Pacific El Niño phases (9, 72). We follow Cai et al. (9) in selecting all models with α at least
575 50% of the observed value, meaning -0.17 or less. Tables S3-S6 show the total and selected
576 realizations for each experiment. We also test the sensitivity of our results to using only one
577 realization from each model (Fig. S13).

578 ENSO amplitude and teleconnections in climate models

581 We define ENSO amplitude as the standard deviation of the quadratically detrended E-
 582 index (9, 43). We calculate each climate model simulation’s amplitude in the historical period,
 583 which we define as 1940-2019 to parallel the observational data, and in the future, which we
 584 define as 2020-2099. The 1940-2019 historical period is chosen so that the historical period is
 585 the same length as the future period.

586 We calculate model-based ENSO teleconnections using the same method as the
 587 observations. We perform this calculation separately for the historical and future periods,
 588 standardizing and linearly detrending each country’s temperature and precipitation time series
 589 independently for each period. This method removes mean shifts due to global warming or low-
 590 frequency variability and allows us to isolate the interannual signal of ENSO.

591

592 Economic damages from changes to ENSO

593 Calculating economic damages from warming-driven ENSO changes requires a
 594 counterfactual world where ENSO evolves without rising temperatures. We calculate the
 595 counterfactual ENSO time series for each simulation by re-scaling its future time series to have
 596 the amplitude that simulation had in the historical period. For example, if E-index amplitude
 597 increases by 20% for a given model realization, we calculate its counterfactual E-index time
 598 series by multiplying its future time series by 0.8 (i.e., $0.8 = 1 - 0.2$). This method preserves the
 599 particular sequence of El Niño and La Niña events in the future, since this sequence is assumed
 600 to be unforced (Fig. S12) but eliminates the forced change in ENSO amplitude.

601 We calculate counterfactual ENSO teleconnections with a similar “delta method.” For
 602 each country in each model, we calculate the change in teleconnection value between the
 603 historical and future simulations. We then add this change to each country’s observed
 604 teleconnection value to implicitly bias-correct the model output. The “counterfactual”
 605 teleconnections are thus equal to the observed values and the “future” teleconnections are the
 606 observed-plus-change values.

607 We then calculate the economic effects of changes to ENSO amplitude by comparing the
 608 future and counterfactual time series and teleconnections from each model. For each year t
 609 between 2020 and 2099, we calculate the growth change from year t to year $t+5$ as the difference
 610 between the future and counterfactual time series and teleconnections:

$$611 \Delta g_{i(t+L)} = [\beta_L E_t^{CF} + \Theta_L E_t^{CF} * \tau_i^{CF}] - [\beta_L E_t^F + \Theta_L E_t^F * \tau_i^F] \quad (4)$$

612

613 Here, E^F refers to the future E-index time series and E^{CF} refers to the counterfactual E-index
 614 time series. Similarly, τ^F refers to future teleconnections and τ^{CF} refers to counterfactual
 615 teleconnections. This calculation yields a growth change time series where each value is the
 616 combined effect of the contemporaneous and lagged effects. We then calculate economic growth
 617 caused by changes in ENSO by subtracting these growth change values from the SSP income
 618 growth projections and integrating growth to calculate income; the new time series represent the
 619 deviations from the SSP baselines caused by changes in ENSO amplitude. Damages are
 620 calculated as the difference between this new time series and the SSP baseline. Details of this
 621 procedure can be found in Burke et al. (13). We perform an analogous calculation using the C-
 622 index time series and teleconnections to calculate C-index damages. We note that this procedure
 623 calculates counterfactual income as accumulated over the entire 21st century, rather than
 624 preceding specific events such as in Fig. 2. This distinction is because these two methods are
 625 aimed at answering different questions. In Fig. 2, we are interested in the effects of specific El
 626

627 Niño events, whereas in Fig. 4, we are interested in the accumulated effect of human-caused
628 changes in ENSO over the 21st century.

629 Finally, given the rebound effects observed after ~10 lags, as well as the large
630 uncertainties in models including longer lags (Fig. S8), we adopt a conservative approach to
631 damage persistence in these calculations. Because we cannot confidently identify permanent
632 effects after 14 years, we allow the growth effect of ENSO to rebound to zero 14 years after the
633 event, meaning that each El Niño affects the global economy for 15 years total (14 lags plus a
634 contemporaneous effect). We do this by applying Eqn. 4 for the first six years (year 0 through
635 year 5) using the coefficients from the main 5-lag model, then allowing the effect to plateau for
636 years 6 through 8, then reversing the coefficients and allowing economies to rebound from years
637 9 through 14. Thus, while we prevent El Niño events from having more than 15 years of an
638 effect, this does not mean that their effect is zero; an affected country has lost substantial
639 economic output during those 15 years that is never recovered. Fig. S16 illustrates this
640 schematically. In a sensitivity analysis, we show results if we assume that damages are
641 permanent and never recovered, a choice which yields substantially greater losses as well as
642 greater uncertainty in those losses (Fig. S13d).

643

644 **Supplementary Text**

645 Regression-based teleconnections

646 Our main analysis uses a correlation coefficient to calculate teleconnections, but we also
647 assess the sensitivity of this choice by using partial regression coefficients instead. Using a
648 regression coefficient leads Peru and Ecuador to be strong outliers from the rest of the
649 distribution (fig. S4e), with values at or above 2. Estimating the growth regression with these
650 values leads to large uncertainties as Peru and Ecuador have an outsized influence on the
651 regression (fig. S4e), so the correlation coefficient is a more stable metric for use in the growth
652 regression. However, we emphasize that the effect of El Niño is still strong and statistically
653 significant when using regression coefficients (Fig. S4e), so our results are not an artifact of the
654 choice to use the correlation coefficient.

655

656 Temperature- or precipitation-based teleconnections

657 Our main analysis defines teleconnections using the combination of temperature and
658 precipitation correlations. We can also define teleconnections solely based on the temperature or
659 precipitation portions of the calculation, similar to previous studies that have focused on
660 temperature to define teleconnections (6, 8). Results for this sensitivity analysis are shown in
661 Fig. S5. The temperature-based estimate is similar to that from both temperature and
662 precipitation, but the effect is weaker with precipitation alone. Our interpretation is that
663 aggregating the data to the monthly time scale and country spatial scale dampens the signal of
664 precipitation more than it does temperature. Consistent with this interpretation, empirical
665 climate-economy studies tend to find little effect of precipitation on country-level growth (13,
666 17).

667

668 Cumulative teleconnections

669 By using the maximum of three-month running means, our main teleconnection analysis
670 focuses on countries' short-term extreme exposure to ENSO rather than capturing cumulative
671 exposure over the entire ENSO life cycle. An alternative teleconnection metric which uses the

672 sum of statistically significant ($p < 0.05$) correlation coefficients across the 15 months for each
673 country yields very similar results, with high correlations between this and our original metric
674 and nearly identical marginal growth effects (fig. S5). This analysis implies that focusing on the
675 few months of maximum exposure is sufficient to capture the effects of ENSO on economies
676 broadly.

677 Heterogeneity in historical teleconnections

678 Our main analysis treats teleconnections as constant in time in the observational period.
679 However, sampling variability and changes in ENSO behavior (among other things) may result
680 in temporal heterogeneity in teleconnections. Fig. S17 shows teleconnections calculated in
681 rolling 30-windows over the historical period. Temporal variation is apparent, at least partly due
682 to the shorter time period used to calculate these teleconnections. However, the distribution of
683 teleconnection values is relatively stable, and the average country experiences temporal variation
684 of only about 13% of its mean value. As such, we use the teleconnection values calculated across
685 the entire time period in our main analysis, though we do allow teleconnections to change with
686 forcing in our climate model analysis.

687 Finally, a key consideration in empirical climate-economy studies is the need to
688 aggregate physical variables to the country scale, which is not a geophysically meaningful scale.
689 To understand the implications of this aggregation, we re-calculate E-index teleconnections at
690 the gridded scale (fig. S17). Teleconnections can vary across grid cells, but the average country
691 only experiences within-country spatial variation of about 11% of its mean teleconnection value
692 (fig. S17). Furthermore, population-weighted country-average grid-cell teleconnection values are
693 similar to the original teleconnection values calculated from country-average temperature and
694 precipitation (fig. S17), implying that subnational spatial variation in ENSO teleconnections does
695 not substantially affect our results.

696 Relationship between our work and recent differences-in-differences literature

697 Our empirical framework is very similar to typical “differences-in-differences” (DID)
698 approaches in economics, involving a treatment variable that varies over time (E and C) and a
699 cross-sectional variable that denotes treatment status (τ). A series of recent papers have
700 illustrated problems with traditional DID approaches, especially when treatment effects are
701 heterogeneous in time and space and treatment timing is staggered (73–75). This type of research
702 design can produce inappropriate comparisons between already treated and newly treated units,
703 resulting in average treatment effect estimates that differ in magnitude and sign from the true
704 effects. While novel estimators have been proposed to avoid these problems (76–78), this
705 literature is still emerging and it is not clear that such estimators are designed for settings with
706 continuous treatments that vary year-to-year and have dynamic effects (79). In lieu of using an
707 alternative estimator, we run several robustness tests to examine the heterogeneity of the effects
708 of ENSO over time and space, which can indicate whether our results are biased by this
709 heterogeneity (80). We estimate the effect in rolling thirty-year windows over the 1960-2019
710 sample period, after dropping individual countries, and after dropping individual years (Fig. S6).
711 In all cases, these estimates are quite similar to our main effect, indicating that unmodeled
712 treatment effect heterogeneity should not pose a threat to our main analysis.

713 Value of climate model selection

717 Our climate model selection criterion preserves the benefit of a multi-model ensemble,
718 allowing us to sample structural uncertainty in model representation of ENSO as well as initial-
719 condition uncertainty, while incorporating information about model skill (81). Treating all
720 simulations in a multi-model ensemble equally has been criticized for assuming that all
721 simulations are independent samples that represent the climate system with equal skill (82),
722 especially since CMIP is an ensemble of opportunity rather than a systematic sampling of
723 uncertainty space. Our consideration of model skill provides an ensemble estimate that is likely
724 more accurate than could be achieved without such consideration. Other methods such as bias
725 correction (83, 84) could also improve ensemble skill, but we use the simpler selection criterion
726 based on α given its consistency with the E- and C-indices and its use in the ENSO modeling
727 community.

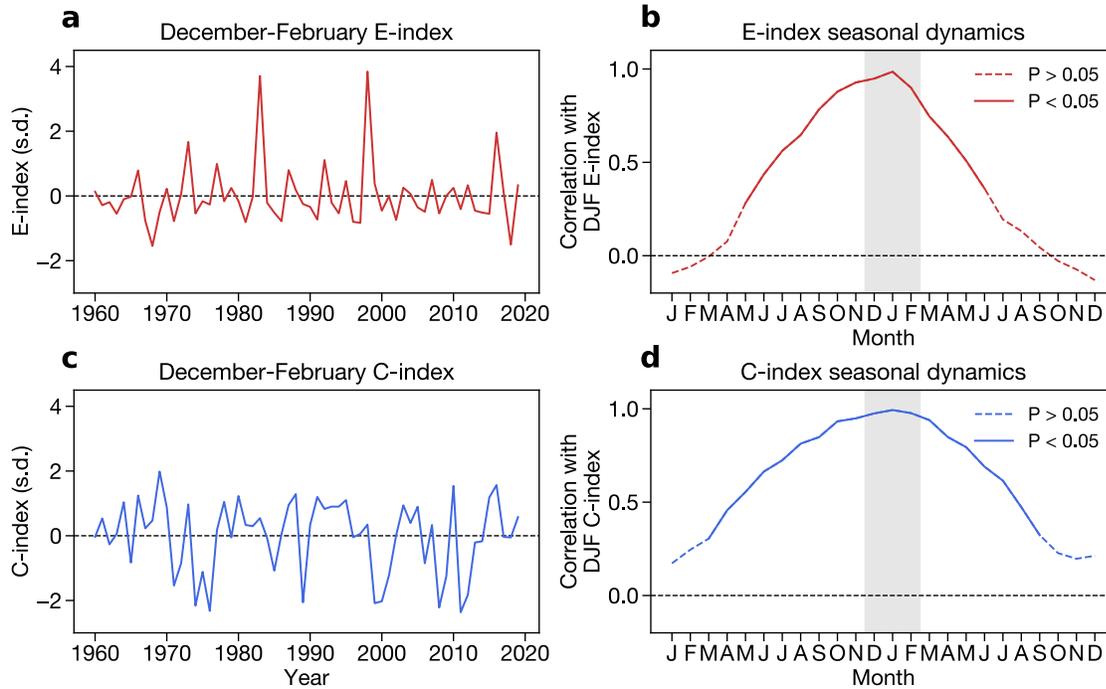
728

729 Sensitivity of damages calculation to alternative choices

730 We incorporate both amplitude and teleconnection changes in our damage projections.
731 Holding teleconnections constant reduces both the magnitude and uncertainty of the damage
732 projections, though they remain negative on average and negatively skewed (fig. S13). Further, a
733 key assumption in these calculations is that the β and θ coefficients (Eqn. 1) remain consistent at
734 a given teleconnection level between the past and future, though individual countries' actual
735 teleconnections may change. This assumption would be violated if societies undertook
736 adaptation measures in response to changes in ENSO amplitude or teleconnections to reduce
737 their sensitivity to ENSO, which is why the need for increased adaptation is a key theme in our
738 results.

739 Finally, our damages calculations use as many simulations from each model as possible
740 (Tables S3-S6) to sample both model structural differences and differences in outcomes due to
741 internal climate variability. Using only the first simulation from each model can generate
742 different results; for example, the SSP5-8.5 simulation yields benefits and SSP1-2.6 yields
743 stronger losses. However, we emphasize that—conditional on our model selection criterion—all
744 selected simulations from a given model are physically plausible given the forcing and boundary
745 conditions. Therefore, the results we present in Fig. 4 are a more complete accounting of the
746 possible effects of ENSO changes.

747

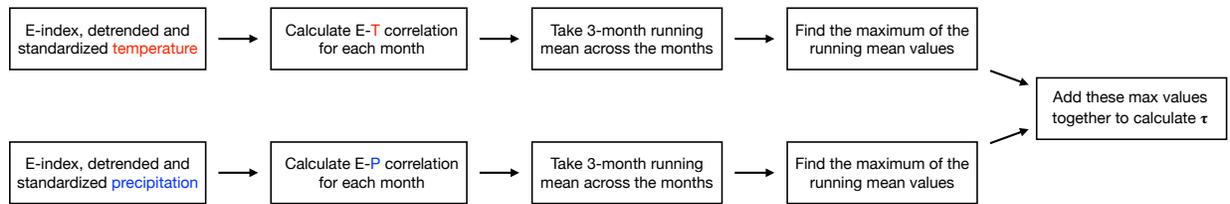


748

749 **Fig. S1.**

750 Interannual and seasonal dynamics of the E- and C-index. A) Timeseries of the average E-index
 751 over December, January, and February (DJF) of each year, where the values are referenced to the
 752 year of January and February. B) Pearson correlation coefficient between the E-index in each
 753 month and the DJF-mean E-index. Solid lines denote correlation coefficients that are statistically
 754 significant ($p < 0.05$) and dashed lines denote correlation coefficients that are statistically
 755 insignificant ($p > 0.05$). C) As in (A), but for the DJF C-index. D) As in (B), but for the DJF C-
 756 index.

757

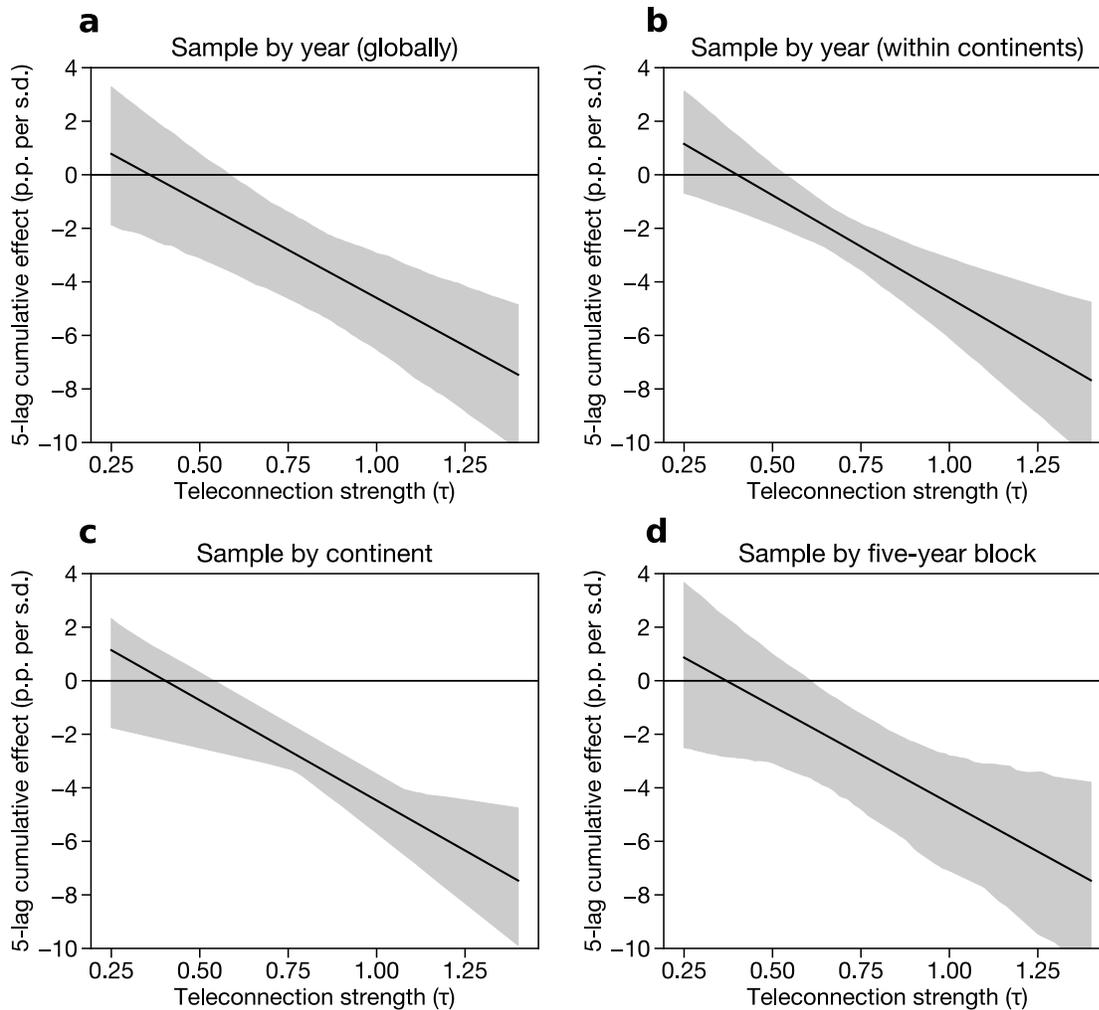


758

759 **Fig. S2**

760 Flow chart for calculation of country-level E-index teleconnections. An analogous calculation is
 761 made for C-index teleconnections.

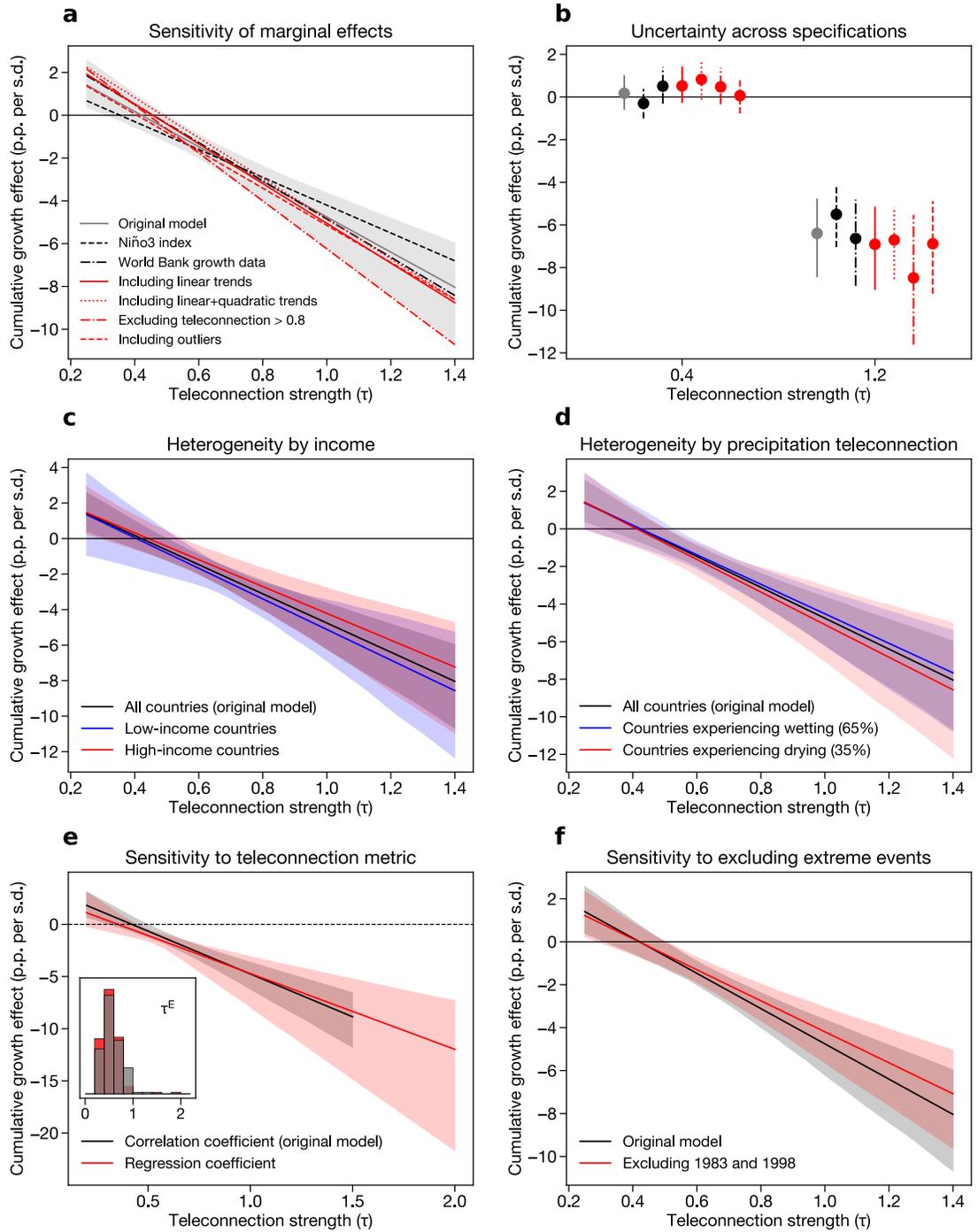
762



763

764 **Fig. S3**

765 Regression results using alternative bootstrap sampling schemes. A) Cumulative 5-lag effect of
 766 ENSO on economic growth when sampling by year, keeping all countries from a given year
 767 together, to account for global spatial correlation in growth within a given year. B) Effect when
 768 sampling by continent-year combinations to account for spatial correlation in growth within
 769 specific continents in a given year. C) Effect when sampling by continents to account for
 770 simultaneous within-continent temporal and spatial correlation in growth. D) Effect when
 771 sampling by five-year blocks to account for global spatial correlation in growth and short-term
 772 (i.e., five-year) temporal correlation in growth. In all cases, solid line shows the mean and
 773 shading shows the 95% confidence intervals. All samples are taken from uniform distributions
 774 with replacement. All axes are the same ranges across panels.

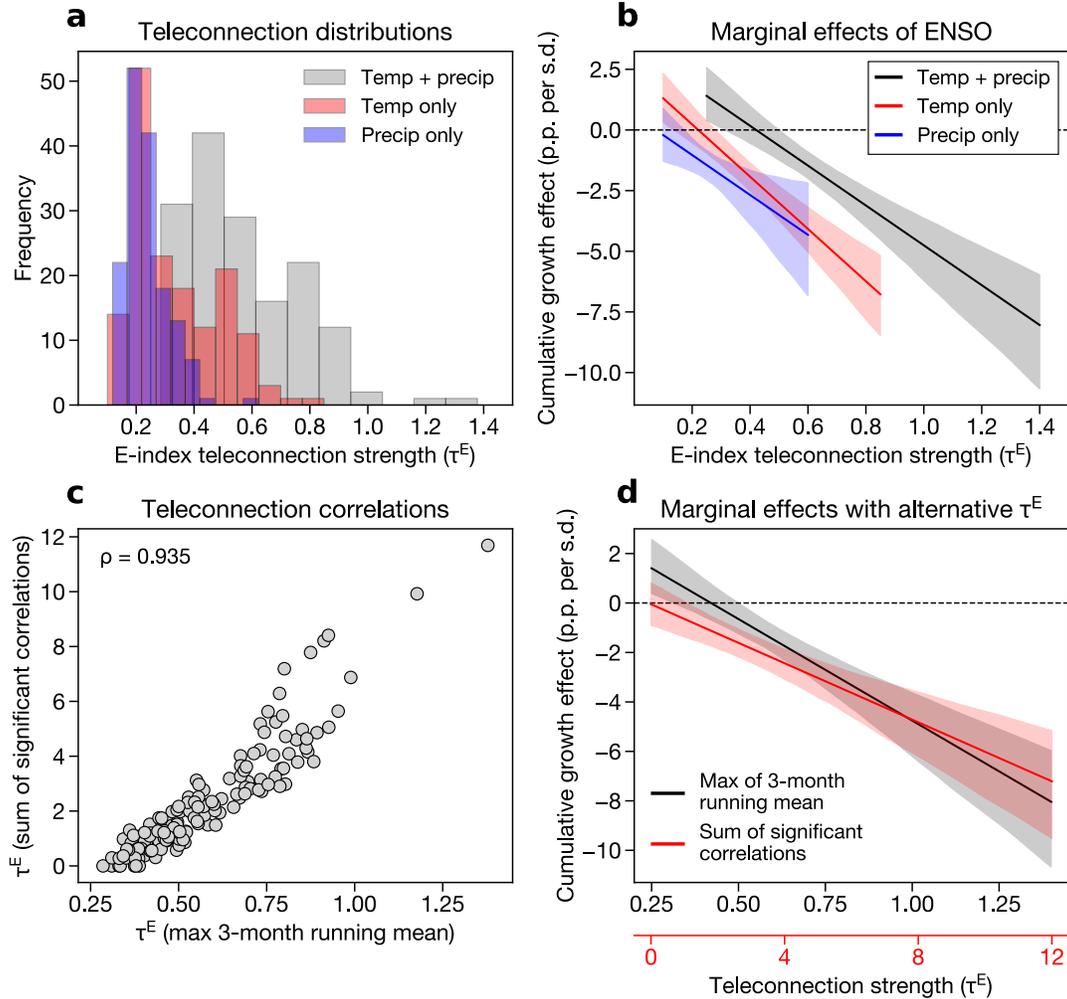


775

776 **Fig. S4**

777 Sensitivity and heterogeneity of the effect of El Niño. A) Cumulative 5-lag effect of El Niño on
 778 growth across a range of specifications: the main model (gray line shows mean and shading
 779 shows 95% confidence intervals), a model using the Niño3 index instead of the E- and C-index
 780 (black dashed line), a model using World Bank growth data instead of the Penn World Tables
 781 (black dash-dot line), a model that includes a country-specific linear trend in growth (red solid
 782 line), a model that includes both linear and quadratic country-specific trends (red dotted line), a

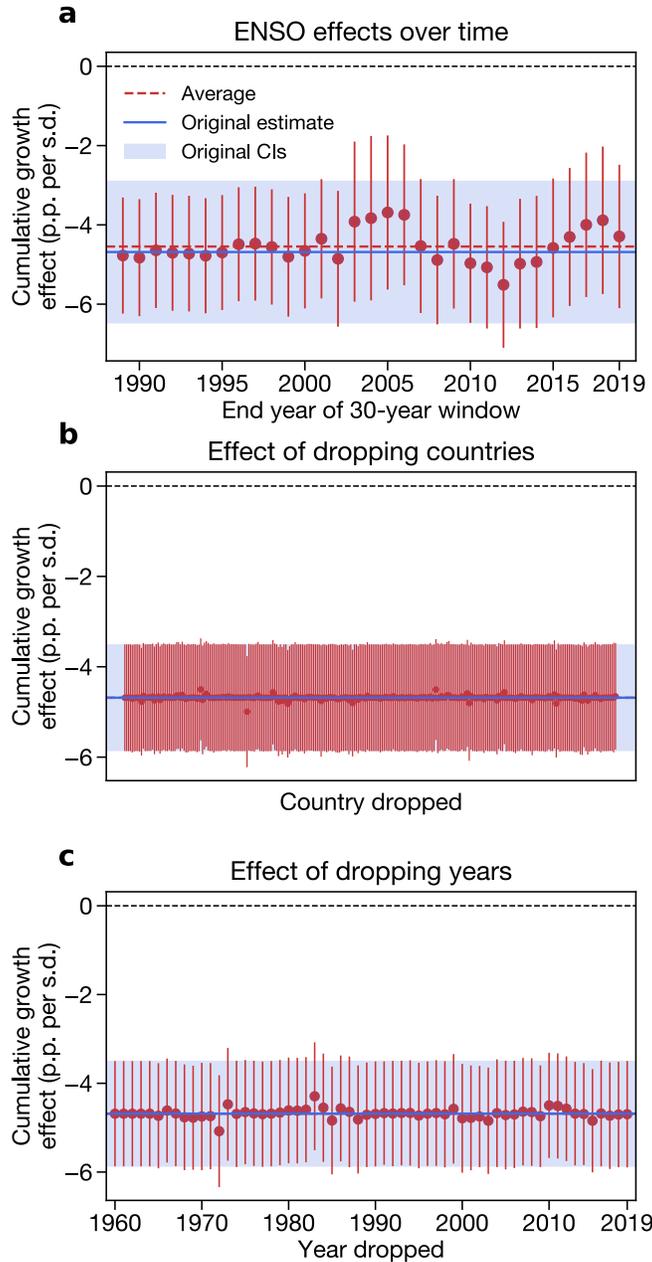
783 model that excludes countries with teleconnection values greater than 0.8 (red dash-dot line), and
784 a model that includes outliers with absolute values of growth greater than 18% (red dashed line).
785 B) Uncertainty in the 5-year cumulative marginal effects of El Niño across each model
786 specification at two representative teleconnection values (0.4 and 1.2). Line styles denote
787 alternative models presented in (A). C) Cumulative marginal effects of El Niño for low-income
788 countries (blue) and high-income countries (red), as defined by World Bank income
789 classifications (Methods). D) Cumulative marginal effects of El Niño for countries experiencing
790 wetting in response to El Niño (positive correlation between the E-index and precipitation, blue)
791 and countries experiencing drying (negative correlation between the E-index and precipitation,
792 red). For each of these samples, we use the original teleconnection value calculated with absolute
793 values in the distributed lag model, but split the sample by the sign of the precipitation
794 teleconnection. In (C) and (D), the original model estimated for all countries is shown in black.
795 E) Cumulative marginal effects of El Niño when using the partial correlation coefficient to
796 measure teleconnections (the main analysis) and when using the regression coefficient instead
797 (red). Inset histograms show the distribution of the two teleconnection metrics. F) Cumulative
798 marginal effects of El Niño when using the full sample (the main analysis, black) and when
799 dropping 1983 and 1998 from the sample (red). In panels (C), (D), (E), and (F), solid line
800 denotes the average and shading denotes 95% confidence intervals from bootstrap resampling by
801 country (Methods).
802
803



804

805 **Fig. S5**

806 Comparison of results using alternative teleconnection metrics. A) Distributions of country-level
 807 teleconnections using monthly temperature correlation coefficients (red), monthly precipitation
 808 correlation coefficients (blue), and their sum (gray). All values are positive since we transform
 809 the correlations to absolute values. B) Cumulative 5-lag effect of ENSO on economic growth
 810 using temperature-only teleconnections (red), precipitation-only teleconnections (blue), and
 811 temperature-plus-precipitation teleconnections (black). C) Relationship between teleconnections
 812 from our main analysis (maximum of three-month running mean) and alternative teleconnections
 813 using the sum of all statistically significant correlation coefficients across the months for each
 814 country. Rho denotes the Spearman's rank correlation coefficient between the two teleconnection
 815 metrics. D) Cumulative 5-lag effect of ENSO on economic growth using the original metric
 816 (black) and the summed correlation coefficient teleconnection metric (red). In (B) and (D), solid
 817 line shows mean and shading shows 95% confidence intervals across 1000 bootstrap iterations,
 818 as in the main analysis.
 819



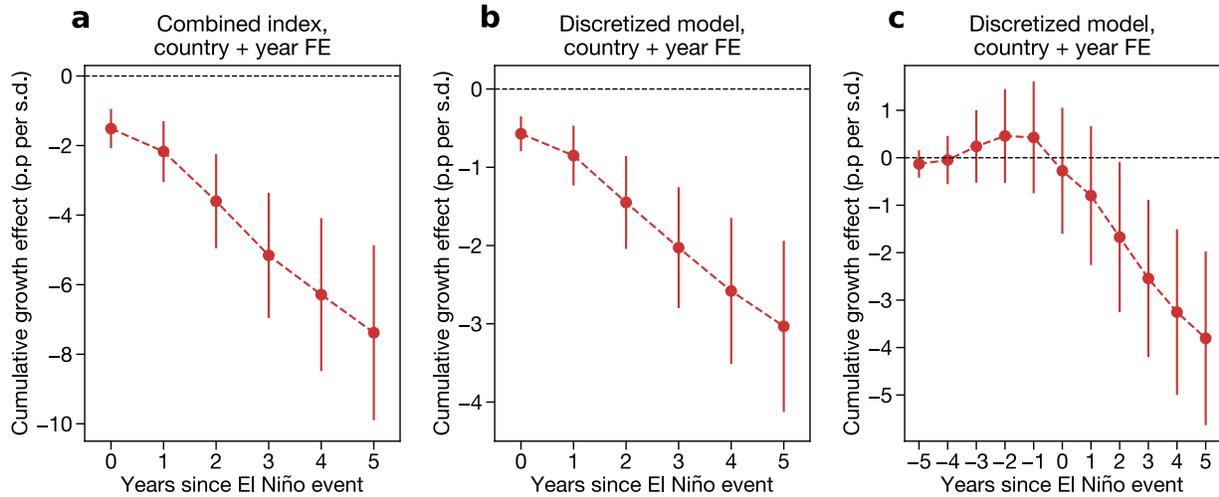
820

821 **Fig. S6**

822 Treatment effect heterogeneity. Panel (A) shows the effect of ENSO on countries with $\tau = 1.0$
 823 calculated in thirty-year rolling windows. X-axis tick refers to the last year of the window. Panel
 824 (B) shows the effect of ENSO on countries with $\tau = 1.0$ when individual countries are dropped
 825 from the sample. We omit country labels for simplicity. Panel (C) shows the effect of ENSO on
 826 countries with $\tau = 1.0$ when individual years are dropped from the sample. In all panels, dashed
 827 red line shows the average effect from all the subsamples, solid blue line shows the central
 828 estimate from our original model, and blue shading shows the 95% confidence interval from our
 829 original model.

830

831



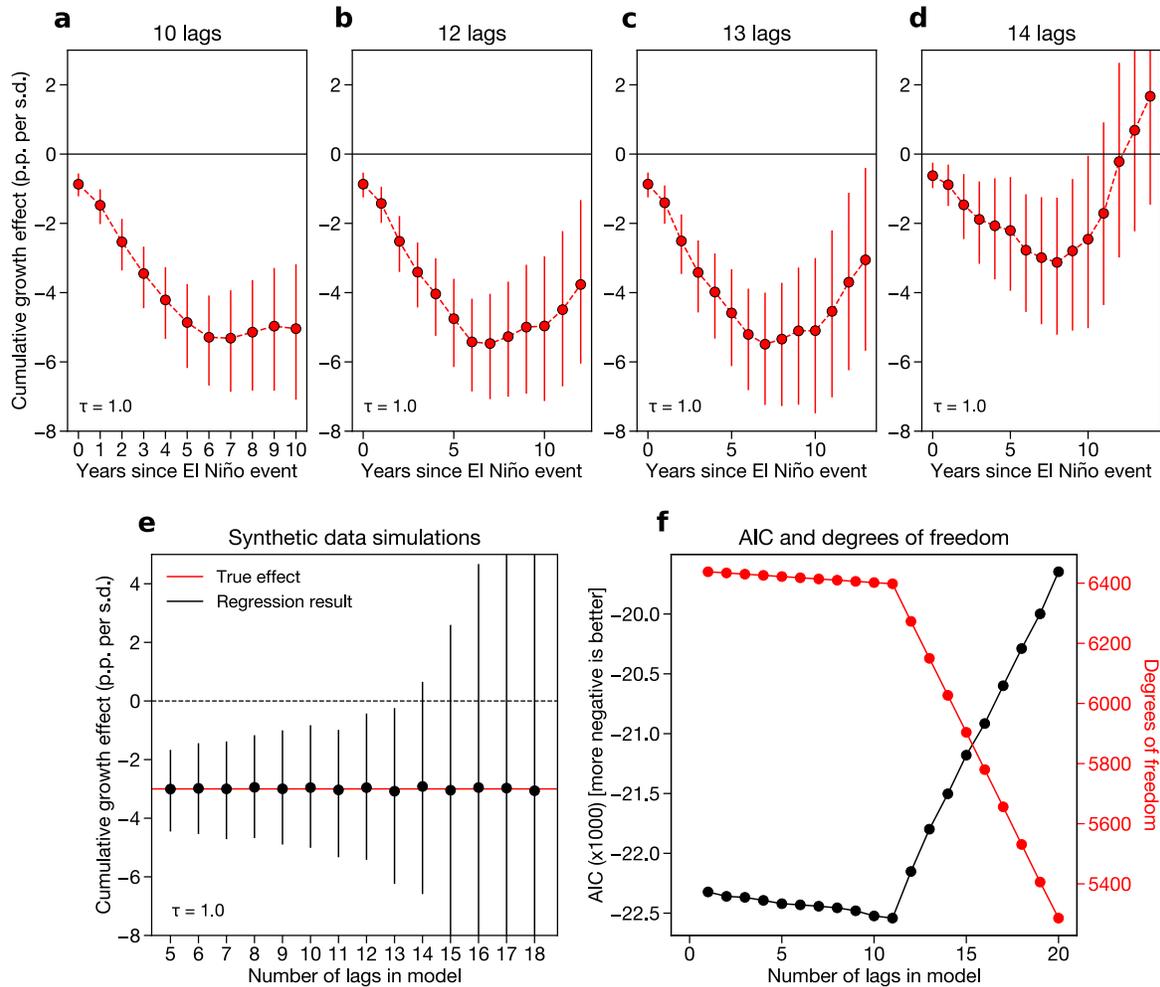
832

833 **Fig. S7**

834 Regression results using several alternative models with both country and year fixed effects.
 835 Panel (A) shows the cumulative effect of a 1-unit increase in the combined index resulting from
 836 multiplying E and τ^E . This index varies in both space and time simultaneously, meaning that both
 837 country and year fixed effects can be included. Panel (B) shows the average cumulative effect of
 838 a 1-s.d. increase in E across all “treated” countries, where treated countries are defined as those
 839 with $\tau^E > 0.5$. Panel (C) shows the same result as (B), with five leads of the E -index added along
 840 with lags. In all cases, the central dashed line shows the mean marginal effect and vertical bars
 841 show the 95% confidence intervals from bootstrap resampling by country.

842

843

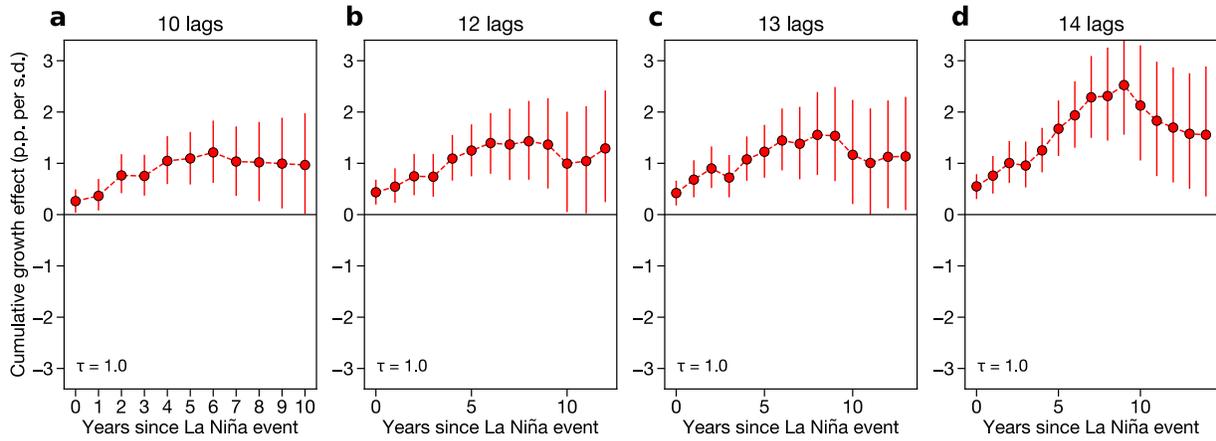


844

845 **Fig. S8**

846 Sensitivity of main regression results to additional lags. A-C) Regression results for countries
 847 with teleconnections greater than or equal to 1.0, estimated with 10 (A), 12 (B), or 15 (C) lags in
 848 the regression model. Confidence intervals are estimated by bootstrap resampling as in the main
 849 analysis. D) Results from synthetic data simulations where a “true” negative ENSO growth effect
 850 is imputed to the data and then estimated using models with lags between 5 and 18 (Methods).
 851 Coefficients estimated using this perfect model framework are shown for a hypothetical country
 852 with $\tau = 1.0$. E) Black line shows Akaike Information Criterion (AIC) values for a series of
 853 regression models with an increasing number of lags from 1 to 20. More negative AIC values are
 854 more desirable. AIC values are divided by 1000 for readability. Red line shows the number of
 855 degrees of freedom for the same set of models.

856

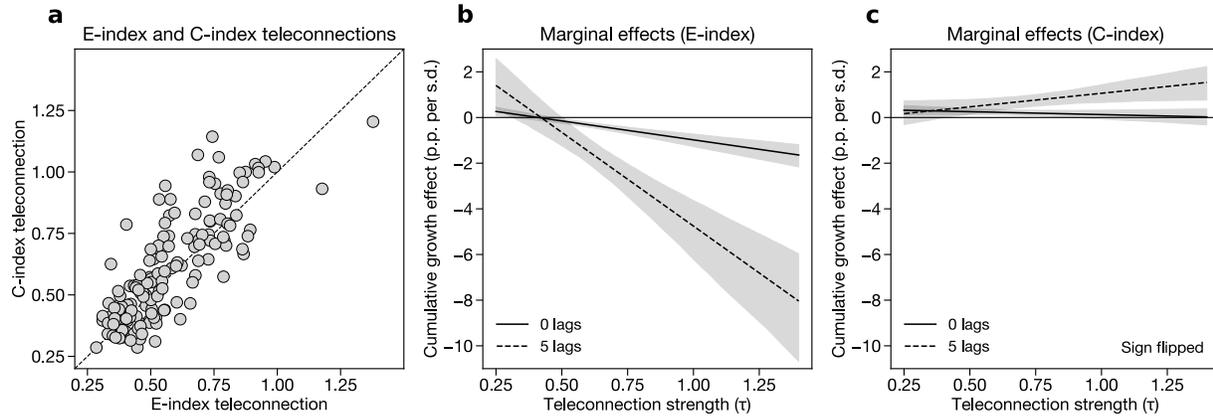


857

858 **Fig. S9**

859 Sensitivity of C-index regression results to additional lags. As in Fig. S7a-d, but for the C-index
 860 coefficients. The sign on the coefficients is flipped to measure the effect of moving from 0 to -1
 861 (i.e., moving a neutral state to a La Niña state).

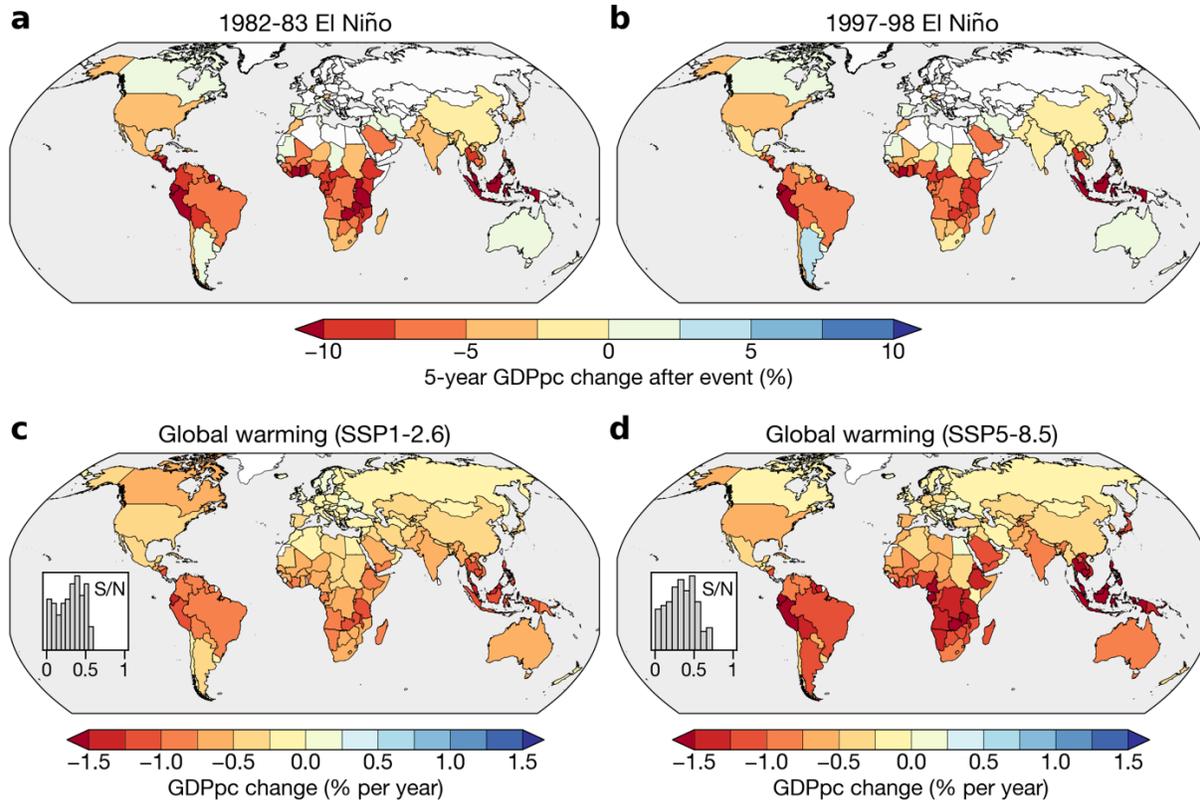
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863

864 **Fig. S10**

865 Teleconnections and marginal effects for both the E-index and C-index. A) Comparison of
 866 country-specific teleconnection metrics calculated using the E-index (x-axis) and C-index (y-
 867 axis). Dashed line denotes the one-to-one line. B) Marginal effects of El Niño (measured by the
 868 E-index) at 0 and 5 lags across a range of teleconnection values. C) Marginal effects of La Niña
 869 (measured by the C-index) at 0 and 5 lags across a range of teleconnection values. The sign on
 870 the coefficients in (C) is flipped to measure the effect of moving from 0 to -1 (i.e., moving from
 871 a neutral state to a La Niña state). In (B) and (C), effects are calculated from a regression that
 872 includes both the E-index and C-index and their corresponding teleconnection metrics
 873 (Methods). Lines denote averages and shading denotes 95% confidence intervals using bootstrap
 874 resampling by country (Methods).

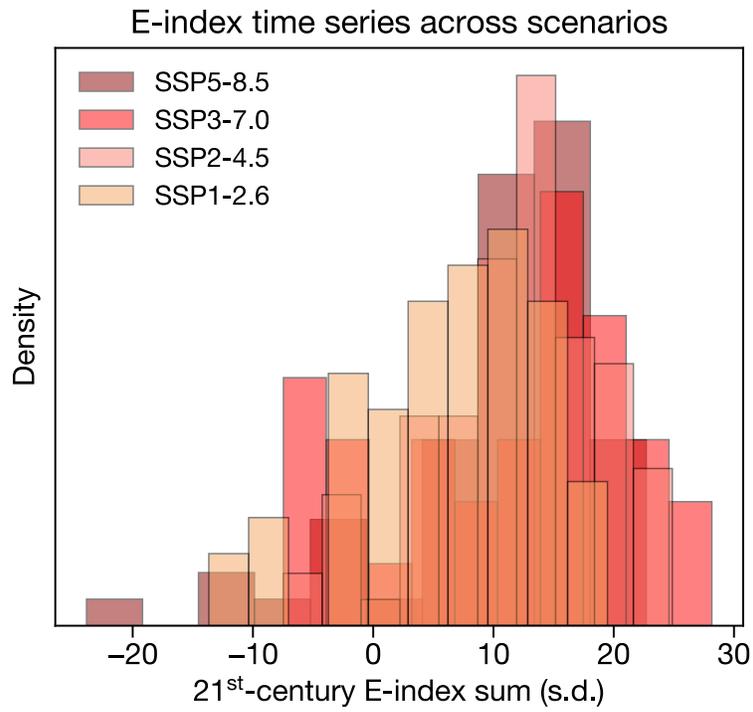


875

876 **Fig. S11**

877 Country-level losses from extreme El Niño events and global warming. A, B) Change in country-
 878 level GDPpc five years after two specific extreme El Niño events: 1982-83 (A) and 1997-98 (B).
 879 Changes are calculated relative to counterfactual trajectories in which the event did not occur
 880 (see Fig. 2a for example of Peru). That is, the color for Brazil in panel B indicates that Brazil's
 881 GDP per capita would have been 5% larger in 2003 if the 1997-98 El Niño event did not occur.
 882 Countries are masked in white if they either have no significant marginal effect of ENSO or do
 883 not have continuous GDPpc data since 1982 (Methods). C, D) 2020-2099 average change in
 884 country-level GDPpc under the SSP1-2.6 (C) and SSP5-8.5 (D) scenarios for the average case
 885 across climate models and regression bootstraps. Insets in C and D show the signal-to-noise
 886 ratios (S/N), meaning the absolute value of the ratio of the ensemble mean GDPpc change to the
 887 ensemble standard deviation GDPpc change. "Ensemble" is defined as all possible combinations
 888 of climate model projections and regression bootstraps.

889

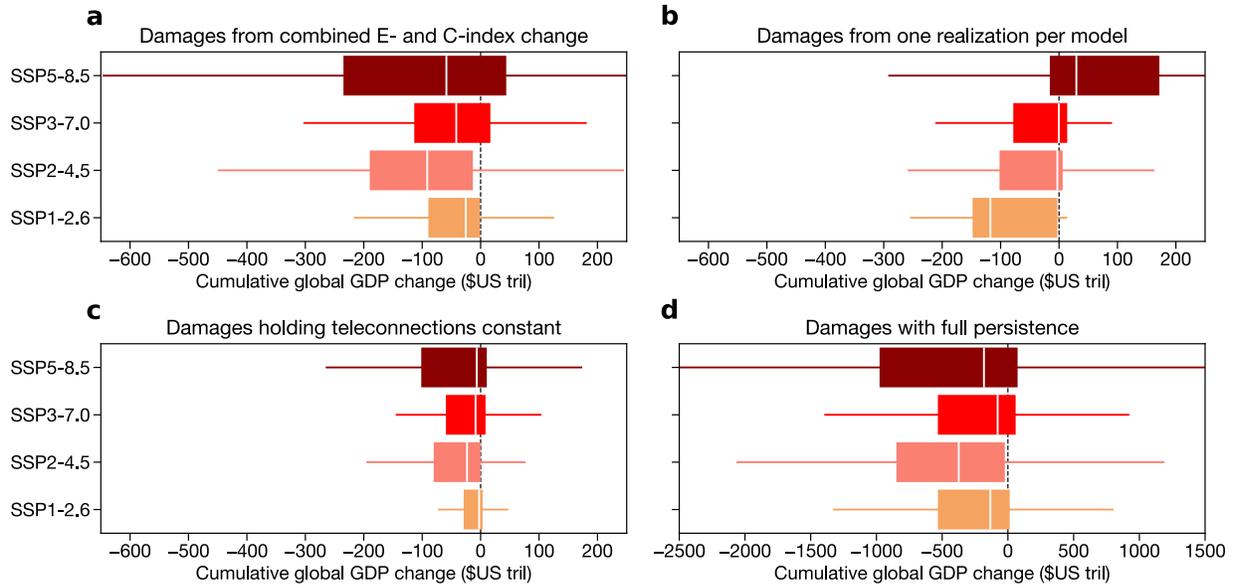


890

891 **Fig. S12**

892 E-index sum across scenarios. Histograms show the distribution of 2020-99 E-index sum values
 893 across simulations within each SSP scenario. Positive values mean that the simulation's E-index
 894 time series has more El Niños than La Niñas.

895



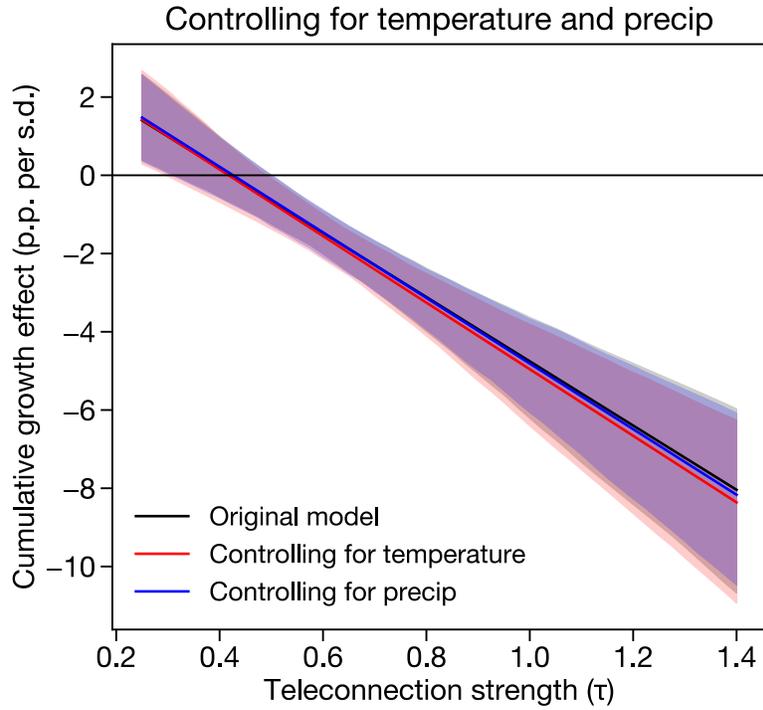
896

897 **Fig. S13**

898 Sensitivity of damage calculations to alternative choices. As in main text Fig. 4a, but for damages due to
 899 the combination of changes in E- and C-index amplitude and teleconnections (A), E-index damages using
 900 only the first realization from each model (B), E-index damages using amplitude change but holding
 901 teleconnections constant (C), and E-index damages when allowing damages to be permanently persistent
 902 (i.e., using the 5-lag model and assuming that the cumulative effects are never recovered) (D). All panels
 903 use a constant 2% discount rate.

904

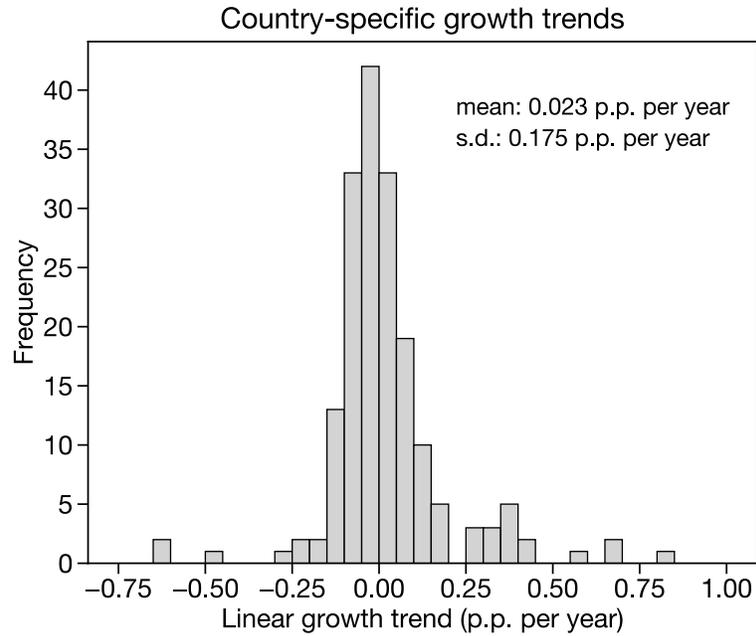
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906

907 **Fig. S14**

908 Effects of controlling for temperature and precipitation in our regression model. Black line
 909 shows results from the original model, red line shows results with the addition of linear and
 910 quadratic terms for country-level annual mean temperature, and blue line shows results with the
 911 addition of linear and quadratic terms for the country-level annual average of monthly total
 912 precipitation.
 913



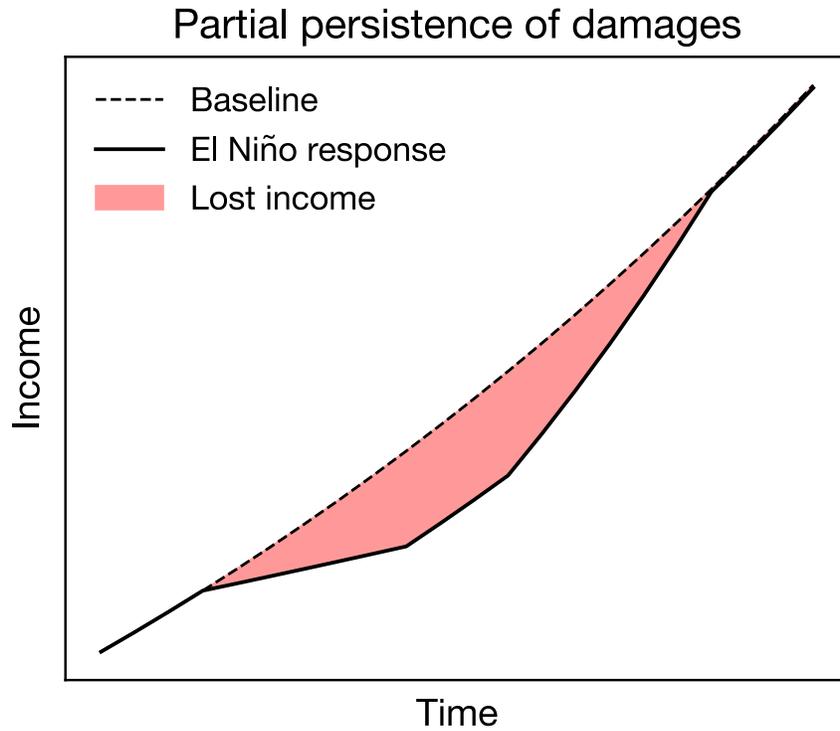
914

915 **Fig. S15**

916 Linear trends in growth. Growth trends are calculated as the linear coefficient on the univariate
 917 regression of each country's growth time series onto time. Only countries with 10 or more years
 918 of growth data are included in this histogram. Text in the top right denotes the mean and standard
 919 deviation of the distribution of trends across countries.

920

921



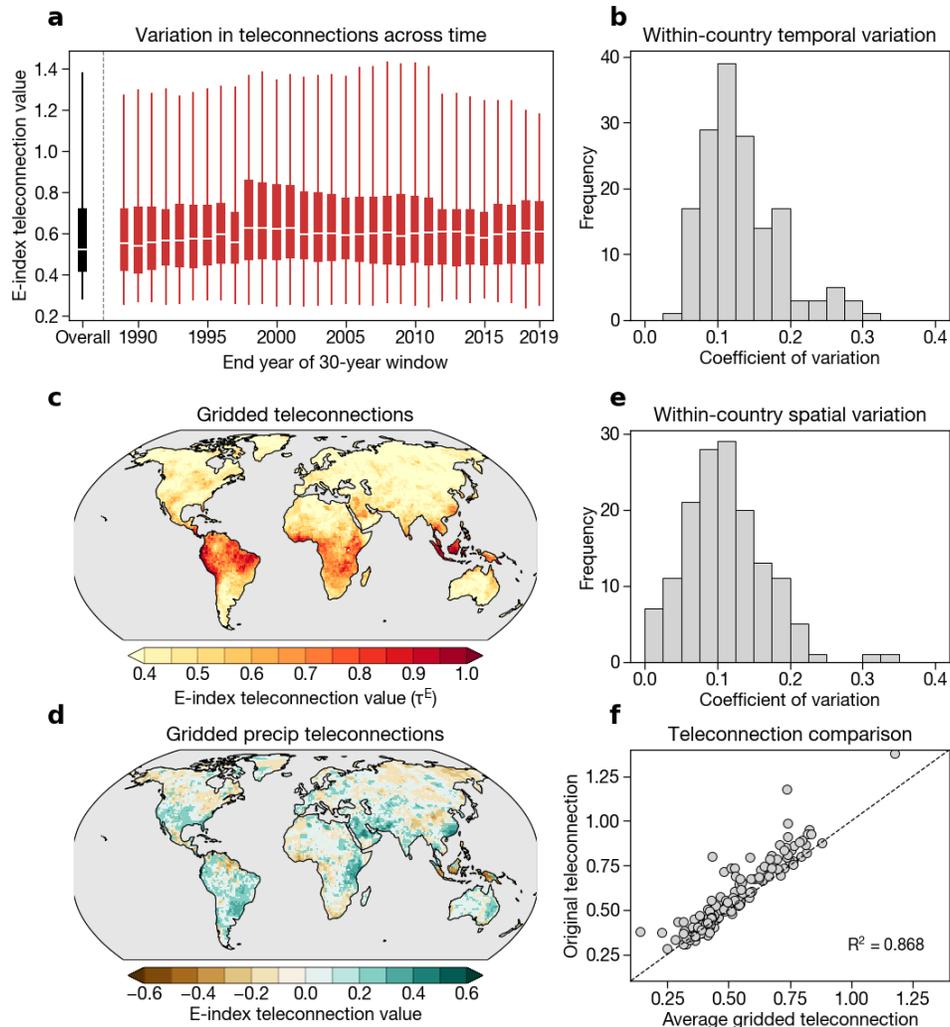
922

923 **Fig. S16**

924 Partial persistence of economic damages. This figure shows a schematic of how we implement
 925 the recovery period in our damage projections. El Niño events negatively affect growth in the
 926 year of the event and in the five years following the event, as in our main model. However, from
 927 years 9 to 14, we allow economies to recover back to their baseline economic trajectory. In the
 928 meantime, there is substantial lost income relative to that baseline trajectory, shown as the red
 929 shaded area.

930

931



932

933 **Fig. S17**

934 Spatiotemporal heterogeneity of observed teleconnections. A) Distribution of E-index
 935 teleconnections in 30-year windows, with x-axis marking the final year of the 30-year window.
 936 An end year of 2015, for example, implies a start year of 1986. The black boxplot shows the
 937 original distribution of teleconnections calculated over the whole 1960-2019 period. White lines
 938 show medians, boxes extend to the 25th and 75th percentiles, and whiskers span the range of the
 939 data. B) Within-country temporal variation, calculated as the coefficient of variation over the 30-
 940 year windows shown in (A). This calculation is performed by dividing the standard deviation of
 941 each country's teleconnection values over all 30-year windows by its mean teleconnection over
 942 those windows. C) Grid-cell E-index teleconnections, calculated using the same method as the
 943 country-level teleconnections, but with standardized grid-cell temperature and precipitation data.
 944 D) Grid-cell precipitation teleconnections, meaning the precipitation component of (C). Note that
 945 the sign is preserved in (D), whereas the teleconnections in (C) and in the main analysis use
 946 absolute values. E) Within-country spatial variation in teleconnections, calculated as the
 947 coefficient of variation of the grid-cell teleconnections when aggregated to the country scale. F)
 948 Relationship between gridded teleconnections averaged at the country scale (with population
 949 weighting) and the original teleconnections using country-average temperature and precipitation.
 950

951 **Table S1.**

952 E-index coefficients with alternative clustering techniques. E-index regression coefficients from
 953 the main regression model (Eqn. 1) using various parametric standard error clustering schemes.
 954 The marginal effect of the E-index for a country i is calculated as the main effect of the E-index
 955 plus the interaction term times $\tau_i^E (\beta + \theta * \tau_i^E, \text{Eqn. 2})$. Clustering accounts for both
 956 spatiotemporal autocorrelation in errors as well as heteroskedasticity across clusters. In all
 957 models, the C-index terms, linear and nonlinear annual mean temperature terms, and the country
 958 fixed effect are included but not shown in the table for simplicity.
 959

	<i>Dependent variable: growth</i>				
	(1)	(2)	(3)	(4)	(5)
$E_t (\beta_0)$	0.0066*** (0.0016)	0.0066* (0.0033)	0.0066 (0.0035)	0.0066* (0.0025)	0.0066 (0.0041)
$E_{t-1} (\beta_1)$	0.0019 (0.0018)	0.0019 (0.0028)	0.0019 (0.0034)	0.0019 (0.0025)	0.0019 (0.0033)
$E_{t-2} (\beta_2)$	0.0054* (0.0022)	0.0054 (0.0044)	0.0054 (0.0043)	0.0054 (0.0041)	0.0054 (0.0039)
$E_{t-3} (\beta_3)$	0.0081*** (0.0021)	0.0081* (0.0036)	0.0081* (0.0035)	0.0081* (0.0029)	0.0081* (0.0032)
$E_{t-4} (\beta_4)$	0.0053** (0.0019)	0.0053 (0.0033)	0.0053* (0.0026)	0.0053* (0.0020)	0.0053* (0.0023)
$E_{t-5} (\beta_5)$	0.0064** (0.0021)	0.0064* (0.0031)	0.0064* (0.0030)	0.0064** (0.0014)	0.0064* (0.0031)
$E_t \times \tau_i^E (\Theta_0)$	-0.0163*** (0.0028)	-0.0163** (0.0055)	-0.0163*** (0.0042)	-0.0163** (0.0036)	-0.0163*** (0.0051)
$E_{t-1} \times \tau_i^E (\Theta_1)$	-0.0072* (0.0028)	-0.0072 (0.0039)	-0.0072 (0.0043)	-0.0072 (0.0030)	-0.0072 (0.0047)
$E_{t-2} \times \tau_i^E (\Theta_2)$	-0.0158*** (0.0036)	-0.0158** (0.0059)	-0.0158*** (0.0048)	-0.0158* (0.0061)	-0.0158*** (0.0046)
$E_{t-3} \times \tau_i^E (\Theta_3)$	-0.0169*** (0.0032)	-0.0169*** (0.0050)	-0.0169*** (0.0042)	-0.0169** (0.0036)	-0.0169*** (0.0038)
$E_{t-4} \times \tau_i^E (\Theta_4)$	-0.0123*** (0.0032)	-0.0123** (0.0045)	-0.0123*** (0.0029)	-0.0123** (0.0024)	-0.0123*** (0.0025)
$E_{t-5} \times \tau_i^E (\Theta_5)$	-0.0121*** (0.0034)	-0.0121* (0.0048)	-0.0121*** (0.0035)	-0.0121*** (0.0010)	-0.0121*** (0.0035)
Observations	7183	7183	7183	7183	7183
Clustering	Country	Year-continent	Year	Continent	Five-year block

*** $p < 0.001$; ** $p < 0.01$; * $p < 0.05$

961 **Table S2.**

962 C-index coefficients with alternative clustering techniques. C-index regression coefficients from
 963 the main regression model (Eqn. 1) using various parametric standard error clustering schemes.
 964 The marginal effect of the C-index for a country i is calculated as the main effect of the C-index
 965 plus the interaction term times τ_i^C ($\phi + \Psi * \tau_i^C$, Eqn. 2). Clustering accounts for both
 966 spatiotemporal autocorrelation in errors as well as heteroskedasticity across clusters. In all
 967 models, the E-index terms, linear and nonlinear annual mean temperature terms, and the country
 968 fixed effect are included but not shown in the table for simplicity.
 969

	<i>Dependent variable: growth</i>				
	(1)	(2)	(3)	(4)	(5)
C_t (Φ_0)	-0.0038* (0.0016)	-0.0038 (0.0028)	-0.0038 (0.0032)	-0.0038 (0.0015)	-0.0038 (0.0052)
C_{t-1} (Φ_1)	0.0048*** (0.0013)	0.0048 (0.0043)	0.0048 (0.0044)	0.0048 (0.0020)	0.0048 (0.0041)
C_{t-2} (Φ_2)	0.0021 (0.0012)	0.0021 (0.0036)	0.0021 (0.0042)	0.0021 (0.0009)	0.0021 (0.0021)
C_{t-3} (Φ_3)	0.0028** (0.0010)	0.0028 (0.0033)	0.0028 (0.0039)	0.0028 (0.0014)	0.0028 (0.0027)
C_{t-4} (Φ_4)	-0.0015 (0.0013)	-0.0015 (0.0035)	-0.0015 (0.0040)	-0.0015 (0.0011)	-0.0015 (0.0021)
C_{t-5} (Φ_5)	-0.0031* (0.0015)	-0.0031 (0.0026)	-0.0031 (0.0032)	-0.0031*** (0.0004)	-0.0031 (0.0042)
$C_t \times \tau_i^C$ (Ψ_0)	0.0026 (0.0023)	0.0026 (0.0039)	0.0026 (0.0042)	0.0026 (0.0017)	0.0026 (0.0067)
$C_{t-1} \times \tau_i^C$ (Ψ_1)	-0.0074*** (0.0018)	-0.0074 (0.0056)	-0.0074 (0.0051)	-0.0074 (0.0031)	-0.0074 (0.0047)
$C_{t-2} \times \tau_i^C$ (Ψ_2)	-0.0059** (0.0018)	-0.0059 (0.0047)	-0.0059 (0.0048)	-0.0059** (0.0011)	-0.0059* (0.0026)
$C_{t-3} \times \tau_i^C$ (Ψ_3)	-0.0041** (0.0015)	-0.0041 (0.0044)	-0.0041 (0.0046)	-0.0041* (0.0013)	-0.0041 (0.0034)
$C_{t-4} \times \tau_i^C$ (Ψ_4)	0.0005 (0.0019)	0.0005 (0.0047)	0.0005 (0.0046)	0.0005 (0.0012)	0.0005 (0.0028)
$C_{t-5} \times \tau_i^C$ (Ψ_5)	0.0025 (0.0020)	0.0025 (0.0035)	0.0025 (0.0034)	0.0025* (0.0007)	0.0025 (0.0045)
Observations	7183	7183	7183	7183	7183
Clustering	Country	Year-continent	Year	Continent	Five-year block

*** $p < 0.001$; ** $p < 0.01$; * $p < 0.05$ 970
971

972 **Table S3.**
 973 CMIP6 models and realizations used from the SSP1-2.6 scenario. Monthly sea surface
 974 temperature (“tos”), monthly atmospheric temperature (“tas”), and daily precipitation (“pr”) are
 975 used from each model. Bolded models are those that have at least 1 realization selected for the
 976 final analysis (Methods).
 977

Model	Total realizations	Selected realizations
CanESM5	50	0
KACE-1-0-G	3	0
MIROC-ES2L	7	7
MIROC6	50	50
MRI-ESM2-0	5	4

978
 979

980 **Table S4.**
 981 CMIP6 models and realizations used from the SSP2-4.5 scenario. Monthly sea surface
 982 temperature (“tos”), monthly atmospheric temperature (“tas”), and daily precipitation (“pr”) are
 983 used from each model. Bolded models are those that have at least 1 realization selected for the
 984 final analysis (Methods).
 985

Model	Total realizations	Selected realizations
ACCESS-CM2	3	0
ACCESS-ESM1-5	11	0
CAMS-CSM1-0	1	0
CESM2	2	0
CESM2-WACCM	3	2
CMCC-CM2-SR5	1	1
CMCC-ESM2	1	1
CNRM-CM6-1	1	0
CanESM5	50	0
EC-Earth3	8	8
FGOALS-g3	3	0
GFDL-ESM4	1	0
HadGEM3-GC31-LL	1	0
INM-CM4-8	1	0
INM-CM5-0	1	0
IPSL-CM6A-LR	5	0
KACE-1-0-G	3	0
MIROC-ES2L	30	30
MIROC6	33	33
MPI-ESM1-2-HR	2	1
MPI-ESM1-2-LR	10	9
NorESM2-LM	2	0
NorESM2-MM	2	1
UKESM1-0-LL	5	0

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 988

989 **Table S5.**
 990 CMIP6 models and realizations used from the SSP3-7.0 scenario. Monthly sea surface
 991 temperature (“tos”), monthly atmospheric temperature (“tas”), and daily precipitation (“pr”) are
 992 used from each model. Bolded models are those that have at least 1 realization selected for the
 993 final analysis (Methods).
 994

Model	Total realizations	Selected realizations
ACCESS-CM2	3	0
ACCESS-ESM1-5	10	0
CAMS-CSM1-0	1	0
CESM2	2	0
CESM2-WACCM	1	1
CMCC-CM2-SR5	1	1
CMCC-ESM2	1	1
CNRM-CM6-1	1	0
CanESM5	50	0
FGOALS-g3	4	0
GFDL-ESM4	1	0
INM-CM4-8	1	0
INM-CM5-0	5	0
IPSL-CM6A-LR	5	0
KACE-1-0-G	3	0
MIROC-ES2L	10	10
MIROC6	3	3
MPI-ESM1-2-HR	10	4
MPI-ESM1-2-LR	7	6
MRI-ESM2-0	5	5
NorESM2-LM	1	1
NorESM2-MM	1	1
UKESM1-0-LL	13	0

995
 996
 997

998 **Table S6.**
 999 CMIP6 models and realizations used from the SSP5-8.5 scenario. Monthly sea surface
 1000 temperature (“tos”), monthly atmospheric temperature (“tas”), and daily precipitation (“pr”) are
 1001 used from each model. Bolded models are those that have at least 1 realization selected for the
 1002 final analysis (Methods).
 1003

Model	Total realizations	Selected realizations
ACCESS-CM2	2	0
ACCESS-ESM1-5	6	0
CAMS-CSM1-0	1	0
CESM2	0	0
CESM2-WACCM	3	1
CMCC-CM2-SR5	1	1
CMCC-ESM2	1	1
CNRM-CM6-1	1	0
CanESM5	50	0
FGOALS-g3	3	0
GFDL-ESM4	1	0
HadGEM3-GC31-LL	4	0
HadGEM3-GC31-MM	4	0
INM-CM4-8	1	0
INM-CM5-0	1	0
IPSL-CM6A-LR	4	0
KACE-1-0-G	3	0
MIROC-ES2L	1	1
MIROC6	50	50
MPI-ESM1-2-HR	2	1
NorESM2-LM	1	1
NorESM2-MM	1	1
UKESM1-0-LL	5	0

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1006 **Table S7.**
 1007 Correlation matrix for the E-index and its lags. Each table entry shows the Pearson correlation
 1008 coefficient between the E-index at various time lags and the E-index at each other time lag.
 1009

	E_t	E_{t-1}	E_{t-2}	E_{t-3}	E_{t-4}	E_{t-5}
E_t		-0.101	-0.335	0.002	0.034	0.002
E_{t-1}			-0.092	-0.336	-0.01	0.037
E_{t-2}				-0.089	-0.291	-0.029
E_{t-3}					-0.094	-0.291
E_{t-4}						-0.076

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1012 **References and Notes**

- 1013 1. M. J. McPhaden, S. E. Zebiak, M. H. Glantz, ENSO as an Integrating Concept in Earth Science.
1014 *Science*. **314**, 1740–1745 (2006).
- 1015 2. J. Bjerknes, ATMOSPHERIC TELECONNECTIONS FROM THE EQUATORIAL PACIFIC. *Mon.*
1016 *Weather Rev.* **97**, 163–172 (1969).
- 1017 3. S.-W. Yeh, W. Cai, S.-K. Min, M. J. McPhaden, D. Dommenget, B. Dewitte, M. Collins, K. Ashok,
1018 S.-I. An, B.-Y. Yim, J.-S. Kug, ENSO Atmospheric Teleconnections and Their Response to
1019 Greenhouse Gas Forcing. *Rev. Geophys.* **56**, 185–206 (2018).
- 1020 4. T. W. Corringham, D. R. Cayan, The Effect of El Niño on Flood Damages in the Western United
1021 States. *Weather Clim. Soc.* **11**, 489–504 (2019).
- 1022 5. F. J. Magilligan, P. S. Goldstein, El Niño floods and culture change: A late Holocene flood history
1023 for the Rio Moquegua, southern Peru. *Geology*. **29**, 431–434 (2001).
- 1024 6. S. M. Hsiang, K. C. Meng, Tropical Economics. *Am. Econ. Rev.* **105**, 257–261 (2015).
- 1025 7. T. Iizumi, J.-J. Luo, A. J. Challinor, G. Sakurai, M. Yokozawa, H. Sakuma, M. E. Brown, T.
1026 Yamagata, Impacts of El Niño Southern Oscillation on the global yields of major crops. *Nat.*
1027 *Commun.* **5**, 3712 (2014).
- 1028 8. S. M. Hsiang, K. C. Meng, M. A. Cane, Civil conflicts are associated with the global climate. *Nature*.
1029 **476**, 438–441 (2011).
- 1030 9. W. Cai, G. Wang, B. Dewitte, L. Wu, A. Santoso, K. Takahashi, Y. Yang, A. Carréric, M. J.
1031 McPhaden, Increased variability of eastern Pacific El Niño under greenhouse warming. *Nature*. **564**,
1032 201–206 (2018).
- 1033 10. W. Cai, B. Ng, G. Wang, A. Santoso, L. Wu, K. Yang, Increased ENSO sea surface temperature
1034 variability under four IPCC emission scenarios. *Nat. Clim. Change*, 1–4 (2022).
- 1035 11. W. Cai, S. Borlace, M. Lengaigne, P. van Rensch, M. Collins, G. Vecchi, A. Timmermann, A.
1036 Santoso, M. J. McPhaden, L. Wu, M. H. England, G. Wang, E. Guilyardi, F.-F. Jin, Increasing
1037 frequency of extreme El Niño events due to greenhouse warming. *Nat. Clim. Change*. **4**, 111–116
1038 (2014).
- 1039 12. W. Cai, A. Santoso, M. Collins, B. Dewitte, C. Karamperidou, J.-S. Kug, M. Lengaigne, M. J.
1040 McPhaden, M. F. Stuecker, A. S. Taschetto, A. Timmermann, L. Wu, S.-W. Yeh, G. Wang, B. Ng,
1041 F. Jia, Y. Yang, J. Ying, X.-T. Zheng, T. Bayr, J. R. Brown, A. Capotondi, K. M. Cobb, B. Gan, T.
1042 Geng, Y.-G. Ham, F.-F. Jin, H.-S. Jo, X. Li, X. Lin, S. McGregor, J.-H. Park, K. Stein, K. Yang, L.
1043 Zhang, W. Zhong, Changing El Niño–Southern Oscillation in a warming climate. *Nat. Rev. Earth*
1044 *Environ.* **2**, 628–644 (2021).
- 1045 13. M. Burke, S. M. Hsiang, E. Miguel, Global non-linear effect of temperature on economic production.
1046 *Nature*. **527**, 235–239 (2015).
- 1047 14. M. Burke, V. Tanutama, Climatic constraints on aggregate economic output. *Natl. Bur. Econ. Res.*
1048 *Work. Pap.* (2019).

- 1049 15. M. Dell, B. F. Jones, B. A. Olken, Temperature shocks and economic growth: Evidence from the last
1050 half century. *Am. Econ. J. Macroecon.* **4**, 66–95 (2012).
- 1051 16. M. Letta, R. S. J. Tol, Weather, Climate and Total Factor Productivity. *Environ. Resour. Econ.* **73**,
1052 283–305 (2019).
- 1053 17. M. Kalkuhl, L. Wenz, The impact of climate conditions on economic production. Evidence from a
1054 global panel of regions. *J. Environ. Econ. Manag.* **103**, 102360 (2020).
- 1055 18. M. Kotz, A. Levermann, L. Wenz, The effect of rainfall changes on economic production. *Nature*.
1056 **601**, 223–227 (2022).
- 1057 19. M. Kotz, L. Wenz, A. Stechemesser, M. Kalkuhl, A. Levermann, Day-to-day temperature variability
1058 reduces economic growth. *Nat. Clim. Change*, 1–7 (2021).
- 1059 20. P. Cashin, K. Mohaddes, M. Raissi, Fair weather or foul? The macroeconomic effects of El Niño. *J.*
1060 *Int. Econ.* **106**, 37–54 (2017).
- 1061 21. S. C. Smith, D. Ubilava, The El Niño Southern Oscillation and Economic Growth in the Developing
1062 World. *Glob. Environ. Change.* **45**, 151–164 (2017).
- 1063 22. R. Generoso, C. Couharde, O. Damette, K. Mohaddes, The Growth Effects of El Niño and La Niña:
1064 Local Weather Conditions Matter. *Ann. Econ. Stat.*, 83–126 (2020).
- 1065 23. A. D. Brunner, El Niño and World Primary Commodity Prices: Warm Water or Hot Air? *Rev. Econ.*
1066 *Stat.* **84**, 176–183 (2002).
- 1067 24. D. Ubilava, El Niño, La Niña, and world coffee price dynamics. *Agric. Econ.* **43**, 17–26 (2012).
- 1068 25. D. Ubilava, The ENSO Effect and Asymmetries in Wheat Price Dynamics. *World Dev.* **96**, 490–502
1069 (2017).
- 1070 26. F. C. Moore, D. B. Diaz, Temperature impacts on economic growth warrant stringent mitigation
1071 policy. *Nat. Clim. Change.* **5**, 127 (2015).
- 1072 27. S. Dietz, N. Stern, Endogenous growth, convexity of damage and climate risk: how Nordhaus’
1073 framework supports deep cuts in carbon emissions. *Econ. J.* **125**, 574–620 (2015).
- 1074 28. E. J. Moyer, M. D. Woolley, N. J. Matteson, M. J. Glotter, D. A. Weisbach, Climate impacts on
1075 economic growth as drivers of uncertainty in the social cost of carbon. *J. Leg. Stud.* **43**, 401–425
1076 (2014).
- 1077 29. K. Takahashi, A. Montecinos, K. Goubanova, B. Dewitte, ENSO regimes: Reinterpreting the
1078 canonical and Modoki El Niño. *Geophys. Res. Lett.* **38** (2011).
- 1079 30. E. M. Rasmusson, T. H. Carpenter, Variations in Tropical Sea Surface Temperature and Surface
1080 Wind Fields Associated with the Southern Oscillation/El Niño. *Mon. Weather Rev.* **110**, 354–384
1081 (1982).
- 1082 31. D. Ubilava, M. Abdolrahimi, The El Niño impact on maize yields is amplified in lower income
1083 teleconnected countries. *Environ. Res. Lett.* **14**, 054008 (2019).

- 1084 32. C. D. Kolstad, F. C. Moore, Estimating the economic impacts of climate change using weather
1085 observations. *Rev. Environ. Econ. Policy*. **14**, 1–24 (2020).
- 1086 33. S. Power, M. Haylock, R. Colman, X. Wang, The Predictability of Interdecadal Changes in ENSO
1087 Activity and ENSO Teleconnections. *J. Clim.* **19**, 4755–4771 (2006).
- 1088 34. N. S. Diffenbaugh, F. V. Davenport, M. Burke, Historical warming has increased U.S. crop insurance
1089 losses. *Environ. Res. Lett.* **16**, 084025 (2021).
- 1090 35. W. Cai, G. Wang, A. Santoso, M. J. McPhaden, L. Wu, F.-F. Jin, A. Timmermann, M. Collins, G.
1091 Vecchi, M. Lengaigne, M. H. England, D. Dommenget, K. Takahashi, E. Guilyardi, Increased
1092 frequency of extreme La Niña events under greenhouse warming. *Nat. Clim. Change*. **5**, 132–137
1093 (2015).
- 1094 36. van Aalst, M.K., Fankhauser, S., Kane, S.M., Sponberg, K., Climate information and forecasting for
1095 development: lessons from the 1997/98 El Niño (2000).
- 1096 37. W. Cai, B. Ng, T. Geng, L. Wu, A. Santoso, M. J. McPhaden, Butterfly effect and a self-modulating
1097 El Niño response to global warming. *Nature*. **585**, 68–73 (2020).
- 1098 38. N. Maher, D. Matei, S. Milinski, J. Marotzke, ENSO change in climate projections: forced response
1099 or internal variability? *Geophys. Res. Lett.* **45**, 11–390 (2018).
- 1100 39. C. Deser, R. Knutti, S. Solomon, A. S. Phillips, Communication of the role of natural variability in
1101 future North American climate. *Nat. Clim. Change*. **2**, 775–779 (2012).
- 1102 40. K.-S. Yun, J.-Y. Lee, A. Timmermann, K. Stein, M. F. Stuecker, J. C. Fyfe, E.-S. Chung, Increasing
1103 ENSO–rainfall variability due to changes in future tropical temperature–rainfall relationship.
1104 *Commun. Earth Environ.* **2**, 1–7 (2021).
- 1105 41. K. Hu, G. Huang, P. Huang, Y. Kosaka, S.-P. Xie, Intensification of El Niño-induced atmospheric
1106 anomalies under greenhouse warming. *Nat. Geosci.* **14**, 377–382 (2021).
- 1107 42. A. T. Wittenberg, A. Rosati, T. L. Delworth, G. A. Vecchi, F. Zeng, ENSO Modulation: Is It
1108 Decadally Predictable? *J. Clim.* **27**, 2667–2681 (2014).
- 1109 43. C. W. Callahan, C. Chen, M. Rugenstein, J. Bloch-Johnson, S. Yang, E. J. Moyer, Robust decrease in
1110 El Niño/Southern Oscillation amplitude under long-term warming. *Nat. Clim. Change*. **11**, 752–757
1111 (2021).
- 1112 44. Council of Economic Advisers, Discounting for public policy: Theory and recent evidence on the
1113 merits of updating the discount rate (2017), (available at
1114 https://obamawhitehouse.archives.gov/sites/default/files/page/files/201701_cea_discounting_issue_brief.pdf).
1115
- 1116 45. M. Meinshausen, J. Lewis, C. McGlade, J. Gütschow, Z. Nicholls, R. Burdon, L. Cozzi, B.
1117 Hackmann, Realization of Paris Agreement pledges may limit warming just below 2 °C. *Nature*.
1118 **604**, 304–309 (2022).
- 1119 46. M. Davis, *Late Victorian Holocausts: El Niño Famines and the Making of the Third World* (Verso
1120 Books, 2002).

- 1121 47. C. W. Callahan, J. S. Mankin, Globally unequal effect of extreme heat on economic growth. *Sci. Adv.*
1122 **8**, eadd3726 (2022).
- 1123 48. B. O'Neill, M. van Aalst, Z. Zaiton Ibrahim, L. Berrang Ford, S. Bhadwal, H. Buhaug, D. Diaz, K.
1124 Frieler, M. Garschagen, A. Magnan, G. Midgely, A. Mirzabaev, A. Thomas, R. Warren, "Key Risks
1125 Across Sectors and Regions" in *Climate Change 2022: Impacts, Adaptation and Vulnerability.*
1126 *Contribution of Working Group II to the Sixth Assessment Report of the Intergovernmental Panel on*
1127 *Climate Change* (Cambridge University Press, Cambridge, UK and New York, NY, USA, 2022),
1128 pp. 2411–2538.
- 1129 49. A. French, R. Mechler, "Managing El Niño risks under uncertainty in Peru" (International Institute
1130 for Applied Systems Analysis, 2017), (available at
1131 [https://pure.iiasa.ac.at/id/eprint/14849/1/French_Mechler_2017_El%20Ni%C3%B1o_Risk_Peru_Re](https://pure.iiasa.ac.at/id/eprint/14849/1/French_Mechler_2017_El%20Ni%C3%B1o_Risk_Peru_Report.pdf)
1132 [port.pdf](https://pure.iiasa.ac.at/id/eprint/14849/1/French_Mechler_2017_El%20Ni%C3%B1o_Risk_Peru_Report.pdf)).
- 1133 50. N. A. Rayner, D. E. Parker, E. B. Horton, C. K. Folland, L. V. Alexander, D. P. Rowell, E. C. Kent,
1134 A. Kaplan, Global analyses of sea surface temperature, sea ice, and night marine air temperature
1135 since the late nineteenth century. *J. Geophys. Res. Atmospheres.* **108** (2003),
1136 doi:10.1029/2002JD002670.
- 1137 51. R. A. Rohde, Z. Hausfather, The Berkeley Earth Land/Ocean Temperature Record. *Earth Syst. Sci.*
1138 *Data Discuss.*, 1–16 (2020).
- 1139 52. U. Schneider, A. Becker, P. Finger, A. Meyer-Christoffer, B. Rudolf, M. Ziese, GPCP full data
1140 reanalysis version 6.0 at 0.5: monthly land-surface precipitation from rain-gauges built on GTS-
1141 based and historic data. *GPCC Data Rep Doi.* **10** (2011).
- 1142 53. C. U. Center for International Earth Science Information Network CIESIN, *Gridded Population of*
1143 *the World, Version 4 (GPWv4): Population Count* (2016).
- 1144 54. R. C. Feenstra, R. Inklaar, M. P. Timmer, The next generation of the Penn World Table. *Am. Econ.*
1145 *Rev.* **105**, 3150–82 (2015).
- 1146 55. A. Deaton, A. Heston, Understanding PPPs and PPP-Based National Accounts. *Am. Econ. J.*
1147 *Macroecon.* **2**, 1–35 (2010).
- 1148 56. T. W. Bank, *World Development Indicators 2016* (2016).
- 1149 57. V. Eyring, S. Bony, G. A. Meehl, C. A. Senior, B. Stevens, R. J. Stouffer, K. E. Taylor, Overview of
1150 the Coupled Model Intercomparison Project Phase 6 (CMIP6) experimental design and
1151 organization. *Geosci. Model Dev.* **9**, 1937–1958 (2016).
- 1152 58. B. C. O'Neill, C. Tebaldi, D. P. van Vuuren, V. Eyring, P. Friedlingstein, G. Hurtt, R. Knutti, E.
1153 Kriegler, J.-F. Lamarque, J. Lowe, G. A. Meehl, R. Moss, K. Riahi, B. M. Sanderson, The Scenario
1154 Model Intercomparison Project (ScenarioMIP) for CMIP6. *Geosci. Model Dev.* **9**, 3461–3482
1155 (2016).
- 1156 59. C. Tebaldi, K. Debeire, V. Eyring, E. Fischer, J. Fyfe, P. Friedlingstein, R. Knutti, J. Lowe, B.
1157 O'Neill, B. Sanderson, D. van Vuuren, K. Riahi, M. Meinshausen, Z. Nicholls, K. B. Tokarska, G.
1158 Hurtt, E. Kriegler, J.-F. Lamarque, G. Meehl, R. Moss, S. E. Bauer, O. Boucher, V. Brovkin, Y.-H.
1159 Byun, M. Dix, S. Gualdi, H. Guo, J. G. John, S. Kharin, Y. Kim, T. Koshiro, L. Ma, D. Olivié, S.

- 1160 Panickal, F. Qiao, X. Rong, N. Rosenbloom, M. Schupfner, R. Séférian, A. Sellar, T. Semmler, X.
 1161 Shi, Z. Song, C. Steger, R. Stouffer, N. Swart, K. Tachiiri, Q. Tang, H. Tatebe, A. Voldoire, E.
 1162 Volodin, K. Wyser, X. Xin, S. Yang, Y. Yu, T. Ziehn, Climate model projections from the Scenario
 1163 Model Intercomparison Project (ScenarioMIP) of CMIP6. *Earth Syst. Dyn.* **12**, 253–293 (2021).
- 1164 60. S. Hoyer, J. Hamman, xarray: N-D labeled arrays and datasets in Python. *J Open Res Softw. Revis.*
 1165 (2017).
- 1166 61. F. Jia, W. Cai, B. Gan, L. Wu, E. Di Lorenzo, Enhanced North Pacific impact on El Niño/Southern
 1167 Oscillation under greenhouse warming. *Nat. Clim. Change*, 1–8 (2021).
- 1168 62. B. Ng, W. Cai, T. Cowan, D. Bi, Impacts of Low-Frequency Internal Climate Variability and
 1169 Greenhouse Warming on El Niño–Southern Oscillation. *J. Clim.* **34**, 2205–2218 (2021).
- 1170 63. E. Tziperman, M. A. Cane, S. E. Zebiak, Y. Xue, B. Blumenthal, Locking of El Niño’s Peak Time to
 1171 the End of the Calendar Year in the Delayed Oscillator Picture of ENSO. *J. Clim.* **11**, 2191–2199
 1172 (1998).
- 1173 64. M. Newman, P. D. Sardeshmukh, Are we near the predictability limit of tropical Indo-Pacific sea
 1174 surface temperatures? *Geophys. Res. Lett.* **44**, 8520–8529 (2017).
- 1175 65. M. Burke, W. M. Davis, N. S. Diffenbaugh, Large potential reduction in economic damages under
 1176 UN mitigation targets. *Nature.* **557** (2018).
- 1177 66. S. M. Hsiang, A. S. Jina, The causal effect of environmental catastrophe on long-run economic
 1178 growth: Evidence from 6,700 cyclones. *Natl. Bur. Econ. Res. Work. Pap.* (2014).
- 1179 67. S. Hsiang, Climate econometrics. *Annu. Rev. Resour. Econ.* **8**, 43–75 (2016).
- 1180 68. J. H. Stock, M. W. Watson, Vector Autoregressions. *J. Econ. Perspect.* **15**, 101–115 (2001).
- 1181 69. N. S. Diffenbaugh, M. Burke, Global warming has increased global economic inequality. *Proc. Natl.*
 1182 *Acad. Sci.* **116**, 9808–9813 (2019).
- 1183 70. E. Guilyardi, A. Wittenberg, A. Fedorov, M. Collins, C. Wang, A. Capotondi, G. J. van Oldenborgh,
 1184 T. Stockdale, Understanding El Niño in Ocean–Atmosphere General Circulation Models: Progress
 1185 and Challenges. *Bull. Am. Meteorol. Soc.* **90**, 325–340 (2009).
- 1186 71. R. Seager, M. Cane, N. Henderson, D.-E. Lee, R. Abernathey, H. Zhang, Strengthening tropical
 1187 Pacific zonal sea surface temperature gradient consistent with rising greenhouse gases. *Nat. Clim.*
 1188 *Change.* **9**, 517–522 (2019).
- 1189 72. C. Karamperidou, F.-F. Jin, J. L. Conroy, The importance of ENSO nonlinearities in tropical pacific
 1190 response to external forcing. *Clim. Dyn.* **49**, 2695–2704 (2017).
- 1191 73. A. C. Baker, D. F. Larcker, C. C. Y. Wang, How much should we trust staggered difference-in-
 1192 differences estimates? *J. Financ. Econ.* **144**, 370–395 (2022).
- 1193 74. C. de Chaisemartin, X. D’Haultfoeuille, Two-Way Fixed Effects and Differences-in-Differences with
 1194 Heterogeneous Treatment Effects: A Survey (2022), , doi:10.3386/w29691.

- 1195 75. J. Roth, P. H. C. Sant'Anna, A. Bilinski, J. Poe, What's Trending in Difference-in-Differences? A
1196 Synthesis of the Recent Econometrics Literature (2023), , doi:10.48550/arXiv.2201.01194.
- 1197 76. B. Callaway, P. H. C. Sant'Anna, Difference-in-Differences with multiple time periods. *J. Econom.*
1198 **225**, 200–230 (2021).
- 1199 77. B. Callaway, A. Goodman-Bacon, P. H. C. Sant'Anna, Difference-in-Differences with a Continuous
1200 Treatment (2021), , doi:10.48550/arXiv.2107.02637.
- 1201 78. C. de Chaisemartin, X. D'Haultfœuille, Two-Way Fixed Effects Estimators with Heterogeneous
1202 Treatment Effects. *Am. Econ. Rev.* **110**, 2964–2996 (2020).
- 1203 79. C. de Chaisemartin, X. D'Haultfoeuille, F. Pasquier, G. Vazquez-Bare, Difference-in-Differences
1204 Estimators for Treatments Continuously Distributed at Every Period (2022), ,
1205 doi:10.48550/arXiv.2201.06898.
- 1206 80. P. Jakiela, Simple Diagnostics for Two-Way Fixed Effects (2021), , doi:10.48550/arXiv.2103.13229.
- 1207 81. R. Knutti, J. Sedláček, B. M. Sanderson, R. Lorenz, E. M. Fischer, V. Eyring, A climate model
1208 projection weighting scheme accounting for performance and interdependence. *Geophys. Res. Lett.*
1209 **44**, 1909–1918 (2017).
- 1210 82. R. Knutti, The end of model democracy? *Clim. Change.* **102**, 395–404 (2010).
- 1211 83. P. Huang, J. Ying, A Multimodel Ensemble Pattern Regression Method to Correct the Tropical
1212 Pacific SST Change Patterns under Global Warming. *J. Clim.* **28**, 4706–4723 (2015).
- 1213 84. D. Chen, M. A. Cane, S. E. Zebiak, R. Cañizares, A. Kaplan, Bias correction of an ocean-atmosphere
1214 coupled model. *Geophys. Res. Lett.* **27**, 2585–2588 (2000).
- 1215