

This pre-print has been submitted to the Geological Society of London Special Publication on Subaqueous Mass Movements and Their Consequences. It has undergone peer review but has not yet been formally accepted for publication and is subject to change during the proof preparation stages.

Title: Lessons learned from monitoring of turbidity currents and guidance for future platform designs

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34

35 **1. Introduction**

36 Reports of sequential seafloor cable breaks at the start of the last century provided the first
37 direct evidence of subaqueous avalanches of sediment called ‘turbidity currents’ (Heezen and
38 Ewing, 1952; Shepard, 1954; Heezen & Ewing, 1955; Heezen et al., 1964; Ryan and Heezen,
39 1965; Piper et al., 1988; Pope et al., 2017). These seafloor-hugging flows were shown to be
40 powerful (reaching up to 20 m/s, sustaining speeds of 3-10 m/s on slopes of less than one
41 degree; Hsu et al., 2008; Carter et al., 2014) and capable of transporting large volumes of sand,
42 mud, organic carbon and nutrients across vast (10s-100s of km) distances (Krause et al., 1970;
43 El Robrini et al., 1985; Piper et al., 1988; Mulder et al., 1997). More than one million km of
44 seafloor cables now connect the world; transmitting more than 98% of all digital data
45 communications, including the internet and financial trading (Burnett and Carter, 2017). We are
46 increasingly reliant on this global network, and on networks of subsea pipelines that support a
47 growing demand for energy (Yergin, 2006; Carter, 2010). It is therefore important to understand
48 the hazards posed to this critical seafloor infrastructure by seafloor mass movements, such as
49 turbidity currents, to inform safe routing, geohazard-tolerant design or mitigation measures
50 where necessary (Bruschi et al., 2006; Randolph and White, 2012; Syanhur and Jaya, 2016;
51 Sequeiros et al., 2019). In addition to being potential geohazards, turbidity currents are also
52 globally important agents of particulate transport. We want to know information such as: i) how
53 they are triggered and linked to onshore sedimentary systems; ii) the frequency at which they
54 recur; iii) how they interact with the seafloor; iv) the physical controls on their run-out; and v)
55 their internal velocity and sediment concentration structure. Inferences can be gleaned from the
56 study of ancient deposits, through analogue modelling of scaled-down flows in the laboratory,
57 and from numerical modelling; however, direct field-scale measurements are needed to calibrate
58 and/or validate all of these approaches (Xu, 2011; Fildani, 2017).

59

60 **1.1. A very brief history of monitoring turbidity currents**

61 Monitoring turbidity currents poses several challenges because deploying instruments on the
62 deep seafloor is logistically challenging, flows may occur infrequently, and the powerful nature
63 of flows can damage the instruments intended to measure them (e.g. Inman et al., 1976; Talling
64 et al., 2013; Puig et al., 2014; Clare et al., 2017; Lintern et al., 2019). Despite these challenges,
65 several studies have prevailed to provide direct measurements of turbidity currents, including
66 seminal field campaigns using point current meters (that measured velocity at one elevation in
67 the water column), in settings ranging from active river-fed fjords (Hay et al., 1982, 1987a&b;

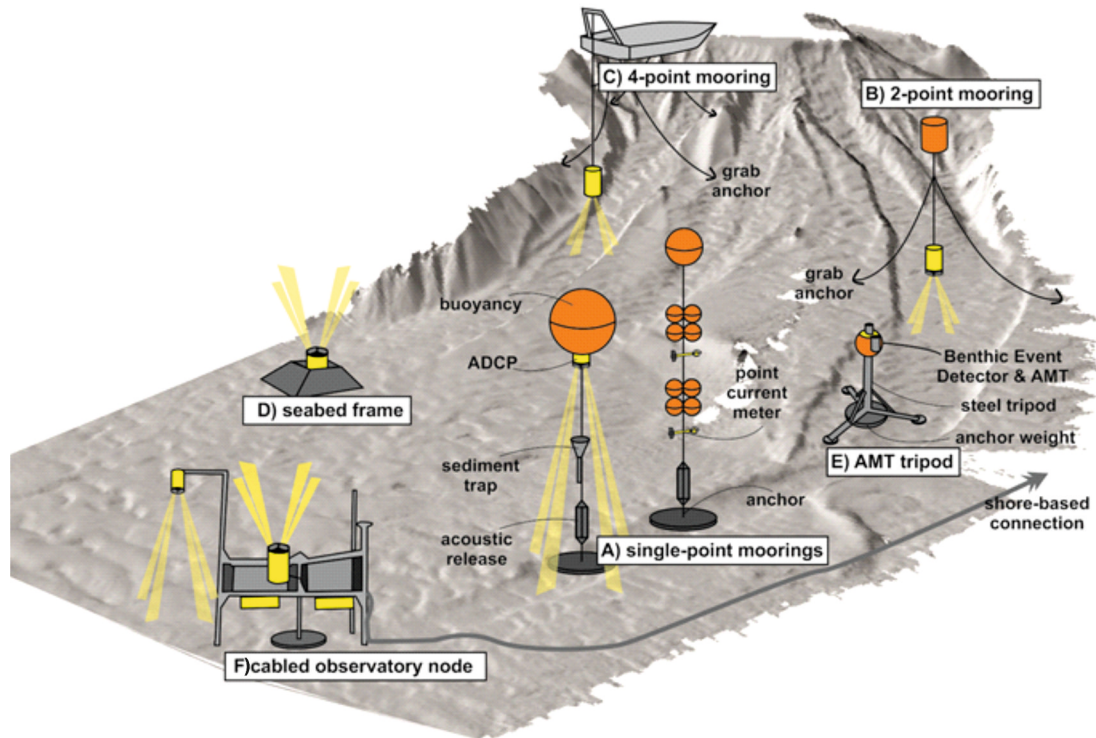
68 Prior et al., 1987; Syvitski and Hein, 1991; Bornhold et al., 1994), lakes (Lambert and
69 Giavanoli, 1988) and deep-sea submarine canyons (Inman et al. 1976; Shepard et al., 1977;
70 Khripounoff et al., 2003, 2009; Vangriesheim et al., 2009). These initial pioneering studies
71 demonstrated that some systems can feature tens of turbidity currents in a year, and that it is
72 feasible to measure flows of up to 3.5 m/s (Prior et al., 1987). These studies were not without
73 incident, however. Many involved damaged or lost instruments (Table 1). Those early studies
74 were also limited with respect to the temporal resolution of measurements, data storage
75 capabilities, duration of deployments, and did not permit depth-resolved flow measurements
76 (Talling et al., 2013).

77 Recent developments in technology, most notably the development of instruments such as
78 Acoustic Doppler Current Profilers (ADCPs) and long-endurance lithium batteries, have
79 enabled depth-resolved measurements of velocity and acoustic backscatter (a proxy
80 measurement for sediment concentration; Thorne and Hanes, 2002) (Cacchione et al., 2006;
81 Shih, 2012). Downward-looking ADCPs avoid the need to place numerous individual point
82 measurements made from within flows (Xu, 2011; Khripounoff et al., 2012). In recent years, a
83 growing number of ADCP-based measurements of turbidity currents have been made in
84 locations including submarine canyons and channels offshore California (Xu et al., 2004; Puig
85 et al., 2004; Xu et al., 2010; Paull et al., 2018), Mississippi (Ross et al., 2009), North-East
86 Atlantic (de Stigter et al., 2007; Martin et al., 2011; Mulder et al., 2012), Mediterranean
87 (Khripounoff et al., 2012; Puig et al., 2012; Martin et al., 2014; Ribó et al., 2015) British
88 Columbia (Hughes Clarke, 2016; Lintern et al., 2016; Hage et al., 2018, 2019), West Africa
89 (Cooper et al., 2013; 2016; Azpiroz-Zabala et al., 2017a&b) and Taiwan (Liu et al., 2012;
90 Zhang et al., 2018).

91

92 Modern turbidity current monitoring campaigns typically integrate multiple sensors and tools,
93 such as multi-beam sonar (imaging the water column), optical back-scatter sensors (to detect
94 suspended particles), acoustic monitoring transponders (to determine seafloor movement),
95 sediment traps (to collect suspended sediment) (Lintern and Hill, 2010; Xu, 2011; Khripounoff
96 et al., 2012; Hughes Clarke, 2016; Lintern et al., 2016; Clare et al., 2017; Paull et al., 2018;
97 Lintern et al., 2019; Hage et al., 2019; Maier et al., 2019a&b). The tools that can be used to
98 measure turbidity currents are partly covered by a number of reviews (Xu, 2011; Talling et al.,
99 2013; Puig et al., 2014; Clare et al., 2017). Here, we focus on the platforms on which these
100 instruments or sensors are mounted, that may include devices such as moorings or frames

101 installed on the seafloor, and may be autonomous or connected via a cabled power and
102 communications link. Examples of different types of platforms are illustrated in Figure 1.
103



104
105 **Figure 1: Illustration depicting examples of some turbidity current monitoring platforms**
106 **discussed in this paper, including: A) Single-point moorings (examples showing older**
107 **point current meters (right) and more recent ADCP designs (left)) with anchors in the**
108 **submarine channel axis; B) Two-point mooring to suspend down-looking instrument**
109 **above active submarine channel , which avoids placement of the anchor in channel axis;**
110 **C) Four-point mooring to stabilise the orientation of a vessel and to enable deployment of**
111 **suspended instruments (Hughes Clarke, 2016); D) Seabed frame to deploy upward-facing**
112 **instrument; E) Acoustic Monitoring Transponder (AMT) tripod with Benthic Event**
113 **Detector (BED) to track movement (Paull et al., 2018); F) Platform connected to a seafloor**
114 **cable network that may host many instruments with real-time communications and power**
115 **(Lintern et al., 2016).**

116

117 1.2. Aims

118 Recent findings enable us to test, refute and refine established hypotheses in turbidity current
119 science; however, direct measurements only exist from a relatively small number of sites

120 worldwide. Many types of system and regions remain completely unrepresented. To date, no
121 detailed measurements of velocity or sediment concentration have been published in water
122 depths of >2 km and none from source to deep-water sink (e.g. submarine fan) as the logistics of
123 placing platforms in deep water remains challenging.

124 Our overarching aim is to share lessons learned from recent campaigns measuring powerful
125 turbidity currents to enable more measurements to be made in a wider variety of locations and
126 settings worldwide. We do this through the following specific objectives. First, we provide an
127 overview of the challenges encountered during the measurement of powerful turbidity currents
128 (up to 10 m/s), including the tilting, displacement and damage of monitoring platforms. We
129 illustrate these challenges with examples from systems including fjord-head deltas, a major
130 river-fed canyon and an oceanographically-fed canyon. Second we introduce single-point
131 moorings and how a successful design for monitoring turbidity currents may differ from that used for
132 more routine oceanographic purposes. These differences include requirements for extra anchor
133 weighting, positive buoyancy, and we discuss the implications of deploying large surface area
134 instruments, such as sediment traps, that can induce excess drag on the mooring string. We
135 outline several methods to reduce drag, and enhance mooring stability. Third, we present a
136 method to deploy two- and four-point moorings, anchored either side of a channel; ensuring that
137 neither the instrument, nor the mooring line, interacts with flows. This is important where
138 pronounced erosion or deposition may occur in the channel axis, and to reduce mooring drag
139 and tilt. Fourth, we assess the deployment of benthic landers and frame-based platforms,
140 describing methods to enhance stability. Finally, we conclude with a discussion on future
141 advances, in both sensor deployment and platform design, which will enable longer endurance
142 turbidity current monitoring.

143

144 **2. Study areas and monitoring data**

145 We now introduce the case study sites discussed in this paper where frequent (sub-annual)
146 turbidity currents have been measured (Figure 2).

147

148 **2.1. Congo Canyon, West Africa**

149 The Congo Canyon is the proximal part of one of the largest submarine channel systems on the
150 planet and is fed directly by the Congo River (Heezen et al., 1964; Babonneau et al., 2010;
151 Azpiroz-Zabala et al., 2017b). Here we focus on previously-published ADCP measurements in
152 the upper part of the Congo Canyon (2 km water depth) that revealed a high frequency of
153 turbidity current activity (Figure 2A; Cooper et al., 2013). Eleven turbidity currents were

154 measured using a downward-looking ADCP (measuring every 5 seconds) deployed from single-
155 point moorings. Flows reached velocities of up to 2.5 m/s and lasted up to 10 days in duration,
156 accounting for 30% of the four-month monitoring period (Azpiroz-Zabala et al., 2017a).

157

158 **2.2. Monterey Canyon, Pacific coast, USA**

159 Monterey Canyon extends from its shelf-incising head in Monterey Bay to the deep-sea
160 Monterey Fan, and is one of the largest submarine canyons on the Pacific Coast of North
161 America (Normark and Carlson, 2003; Paull et al., 2005). Sediment is supplied to the canyon
162 head by long-shore sediment transport cells, rather than directly from a river source (Best and
163 Griggs, 1991). Frequent turbidity currents have been recorded by numerous studies in the
164 canyon using downward-looking ADCPs on single-point moorings (e.g. Xu and Noble 2009;
165 Xu et al., 2013; 2014). A recent (2015-2017) 18-month coordinated international experiment
166 installed more than 50 sensors within the canyon to record the passage of 15 turbidity currents;
167 some of which ran out for >50 km in water depths of up to 1840 m and reached velocities of
168 >7.2 m/s (Paull et al., 2018; Figure 2B). Here, we focus four different types of platform: i) a
169 downward-looking ADCP and sediment trap (at 290 m water depth; Maier et al., 2019a); ii) a
170 800 kg tripod frame (deployed at 300 m water depth) fitted with an Acoustic Monitoring
171 Transponder (AMT) and Benthic Event Detector (BED) to track its movement (Paull et al.,
172 2018; Urlaub et al., 2018); and iii) a seafloor frame deployed at the distal end of the monitoring
173 array (1840 m water depth) that hosted numerous instruments including upward-looking ADCPs
174 (Paull et al., 2018).

175

176 **2.3. Squamish prodelta, Canadian Pacific Coast**

177 The Squamish prodelta lies offshore from the Squamish River that drains into the Howe Sound
178 fjord, British Columbia. Three submarine channels connect the delta lip to channel lobes in
179 water depths of up to 200 m (Figure 2C: Hughes Clarke, 2016). Repeat seafloor surveys, and
180 water column monitoring has revealed extremely frequent (>100/year) turbidity currents during
181 seasonal peaks in meltwater discharge (Hughes Clarke et al., 2012; Clare et al., 2016). Here we
182 focus on a seafloor frame containing and upward-looking ADCP (installed on the terminal lobe
183 of one of the channels in 2011; Figure 2C), and multi-point moorings installed in 2013 and 2015
184 to measure flows that attained velocities of up to 3 m/s (Hughes Clarke, 2016; Hage et al.,
185 2018).

186

187 **2.4. Bute Inlet, Canadian Pacific Coast**

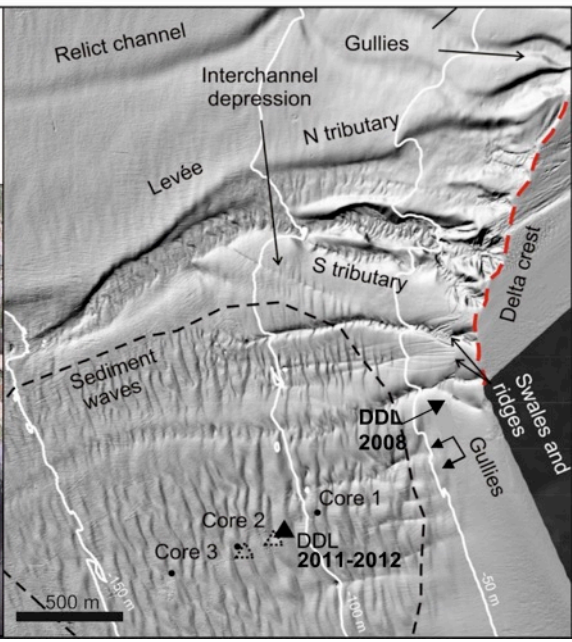
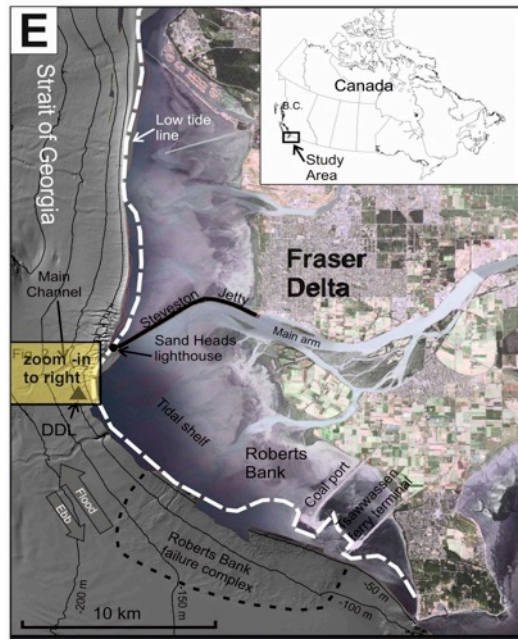
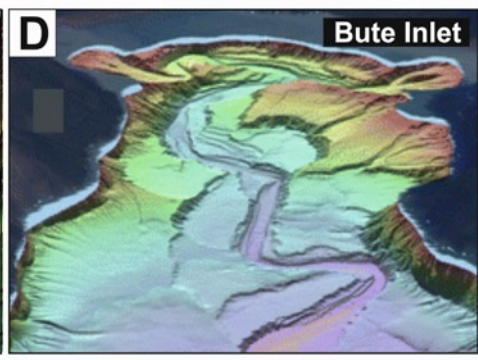
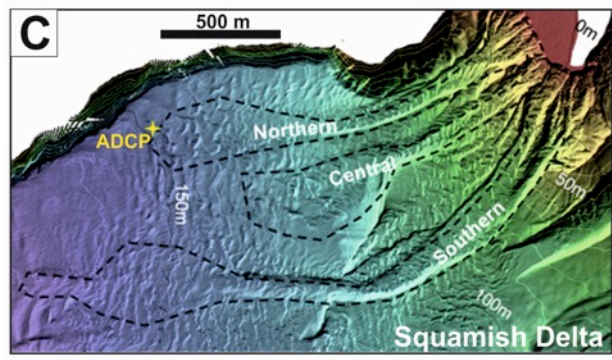
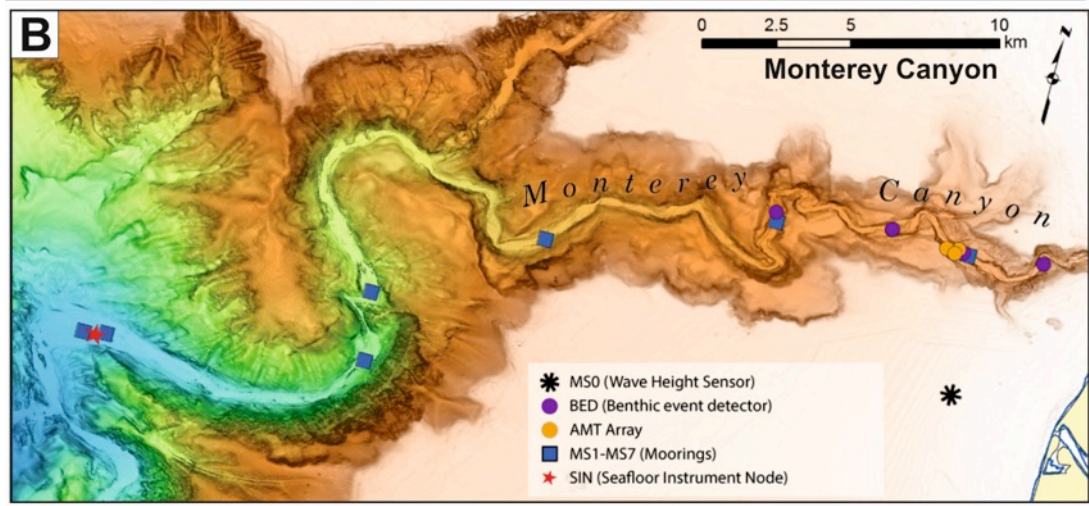
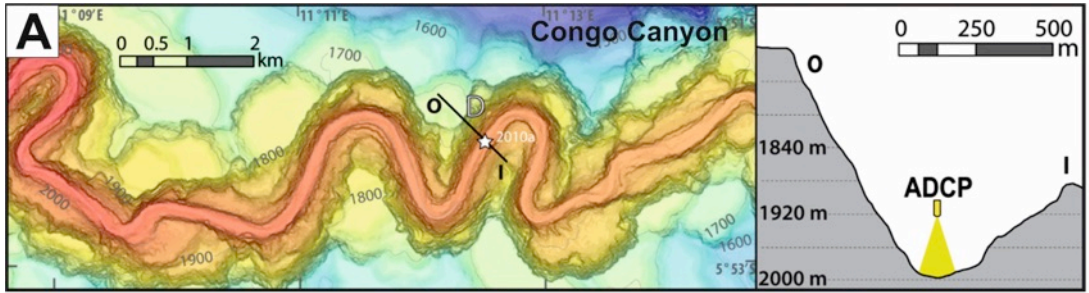
188 Bute Inlet fjord (also in British Columbia) is fed by the Homathko and Southgate rivers, which
189 in turn feed the submarine deltas at the head of a sinuous 50 km-long submarine channel that
190 extends to a terminal lobe at ~700 m water depth (Figure 2D; Prior et al., 1987). Repeated
191 seafloor surveys have shown >metre-scale elevation changes in the channel axis due to erosion
192 and deposition caused by turbidity currents (Gales et al., 2018). Some of the earliest direct
193 measurements of turbidity currents were made in Bute Inlet using point current meters on
194 moorings that recorded flows in excess of 3 m/s (Prior et al., 1987; Zeng et al., 1991). Here, we
195 focus on more recent ADCP- and 500 kHz multibeam echosounder-based measurements of
196 flows using two- and four-point moorings, deployed in 2016 and 2018.

197

198 **2.5. Fraser Delta, Canadian Pacific Coast**

199 The Fraser submarine delta lies offshore from the Fraser River, British Columbia. The principal
200 offshore distributary channel is located immediately seaward of the river outflow, and is flanked
201 to its south by a field of sediment waves on the delta slope (Figure 2E; Lintern et al., 2016).
202 Historical slope failures have been observed from repeat seafloor surveys on the submarine
203 delta slope (e.g. Kostachuk et al., 1992; Hill, 2012). Unlike the previous examples, here we
204 focus on an array of monitoring platforms installed outside of a submarine channel the Delta
205 Dynamics Laboratory (DDL), sited on the open sediment wave field (Figure 2E). The DDL is
206 part of Ocean Network Canada's VENUS cabled network and has been in operation since 2008
207 (Lintern & Hill, 2010; Lintern et al., 2016). The platform can host a wide range of
208 instrumentation due to its cabled power and communications connection, some of which include
209 upward- and downward-looking ADCPs, velocity profilers, turbidity sensors and video camera
210 (Lintern et al., 2016). Other platforms at the site include a seismic liquefaction in situ
211 penetrometer (SLIP), which is measuring pressures and movement within the bed, and a
212 hydrophone array, which is listening for landslides and other noises. As with the Bute and
213 Squamish sites, turbidity currents are frequent during the spring and summer when river
214 discharge is elevated (Ayranci et al., 2012).

215



217 **Figure 2: Location maps and bathymetry for each of the sites discussed in this paper. A:**
218 **Location of ADCP mooring in Congo Canyon, West Africa at 2000 m water depth**
219 **(Modified from Azpiroz-Zabala et al., 2017a). B: Configuration of Monterey Canyon CCE**
220 **instrument deployment, offshore Moss Landing, California, USA. Water depth range of**
221 **instrument deployment was 30 m to 1840 m (from [https://www.mbari.org/cce-](https://www.mbari.org/cce-instruments-2019/)**
222 **[instruments-2019/](https://www.mbari.org/cce-instruments-2019/)).** C: Squamish submarine delta in Howe Sound, British Columbia.
223 **Water depth is up to 200 m (modified from Clare et al., 2016). D: Bute Inlet, British**
224 **Columbia, with water depths of up to 700 m. E: Fraser Delta, British Columbia, showing**
225 **relationship with the Fraser River (left) and detail on offshore delta channel and bedform**
226 **field (right) where the Delta Dynamics Laboratory (DDL) was deployed in different**
227 **locations (modified from Lintern et al., 2016).**

228

229 **3. Results from recent direct monitoring of turbidity currents**

230 We now summarise issues we encountered during recent turbidity current monitoring
231 campaigns, ordered from smallest to greatest impact.

232

233 **3.1. Temporary instability of single-point moorings: pull down, pitch, roll and rotation**

234 Single-point ADCP moorings in a submarine canyon or channel axis commonly record an
235 abrupt increase in water pressure coincident with the arrival of a turbidity current. In the 2015-
236 2017 Monterey Canyon Coordinated Canyon Experiment (CCE), each of the 15 turbidity
237 currents caused an initial increase in water pressure that generally declined over 4 to 120
238 minutes (Paull et al., 2018). This increase in water pressure is attributed to pull-down of the
239 mooring cable, due to drag imparted by the flow front (which reached velocities of up to 7.2
240 m/s) most likely exerted on instruments that were within the flow. A decrease in water pressure
241 occurred when the flows decelerated and the mooring gradually returned to its original vertical
242 position. A similar situation was observed in a previous experiment in Monterey Canyon, where
243 a mooring was severely tilted during the first 15 minutes of a turbidity current, causing a
244 sediment trap (located at 70 m above seafloor) to be pulled down by 37 m into the lower parts of
245 the flow; thus explaining the anomalously coarse material collected by the sediment trap
246 (Symons et al., 2017). Mooring tilt and down-canyon transport also occurred during strong
247 internal tidal flows in Monterey Canyon (i.e. tidal frequency flows trapped within the canyon
248 topography, unrelated to turbidity currents). On November 30th 2015, during a particularly

249 strong up-canyon internal tide (~ 1 m/s) the lower current meter was pulled down 2 m and tilted
250 more than 20 degrees.

251 Such pull-down effects were not observed in the Congo Canyon, where the mooring
252 construction was much simpler and acoustic release links were located much higher above the
253 seafloor than in the Monterey Canyon experiments (Figure 9). This is not to say that the Congo
254 Canyon mooring remained unaffected by flows, however. Intervals of increased pitch, roll and
255 tilting (< 2 degrees) were recorded by the downward-looking ADCP during turbidity currents;
256 dominantly during the initial passage (< 1 hour) of the fast frontal cell. These effects (in
257 particular the rotation of the buoy housing the ADCP), resulted in transient interaction of the
258 ADCP beams with the narrow canyon sidewalls, thus limiting the depth range and quality of
259 velocity and backscatter measurements.

260

261 **3.2. Down-canyon transport of single-point moorings and damage to instruments**

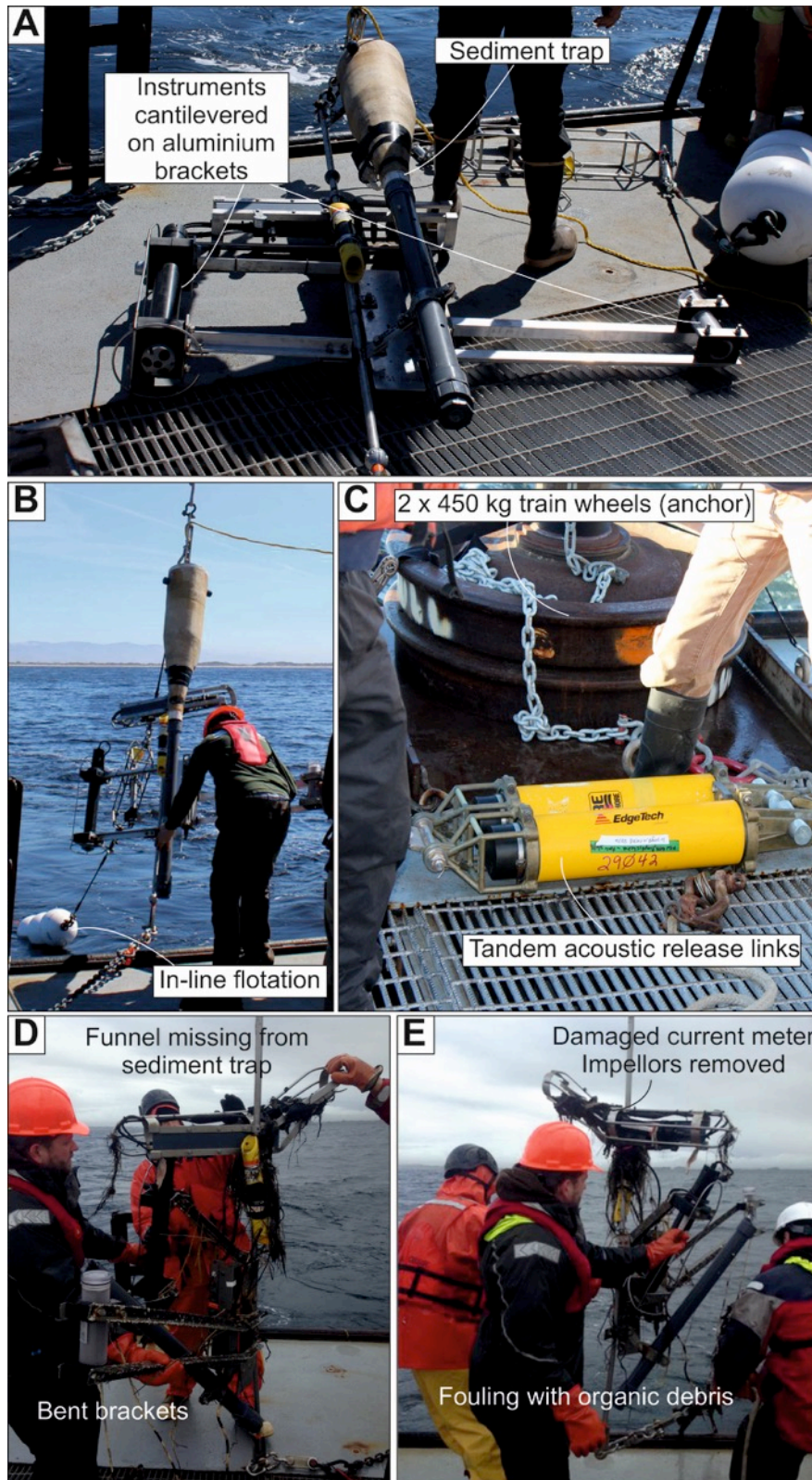
262 As well as the reversible pressure changes noted at the start of turbidity currents, several
263 turbidity currents in Monterey Canyon caused permanent pressure and temperature changes, as
264 recorded by ADCPs on single-point moorings. These irreversible changes indicate that, in
265 addition to the buoy-mounted ADCP being temporarily pulled towards the seafloor, single-point
266 moorings were also transported down-canyon. Symons et al. (2017) documented the 580 m
267 down-canyon transport of a single-point mooring attached to a 1000 kg anchor at a speed of
268 ~ 0.5 m/s from a 2002-2003 deployment (Xu et al., 2004; 2014). During the CCE (December 1st
269 2015), a single-point mooring (using a 450 kg train wheel for an anchor) was moved down
270 canyon (as evidenced by an average drop in pressure of 3 m) by a relatively small turbidity
271 current (~ 3 m/s). The most powerful flow event (January 15th 2016) caused down-canyon
272 transport of the same mooring by 7.1 km, at an average speed of 4.5 m/s (Paull et al., 2018).
273 This mooring ultimately broke loose from its anchor and was retrieved at the sea surface.

274 On the final of three deployments in the Monterey Canyon CCE, two train wheels (~ 900 kg)
275 were used to anchor the single-point mooring and in-line flotation was placed above each
276 sediment trap (as well as additional flotation at the top of the mooring; Figure 3). Mooring
277 performance was much improved by this revised design. Even in very strong turbidity currents
278 (> 5 m/s) the mooring did not move. Tilt and down-pull during strong internal tides were also
279 considerably reduced (< 10 degree and < 1 m, respectively). To make additional measurements
280 within turbidity currents, several instruments were installed on the mooring line beneath the
281 ADCP for the Monterey CCE, including Anderson-style sediment traps, altimeters and point
282 current meters (Figure 3A,B&C). Significant damage was recorded upon retrieval of these

283 instruments, however, including loss of the impellers for the current meter, fouling of
284 instruments with sediment and organic debris, removal of the sediment trap inlet funnel, and
285 sand-blasting, bending and buckling of steel instrument frames (Paull et al., 2018; Maier et al.,
286 2019a; Figure 3D&E). One particularly important issue also concerned damage to the acoustic
287 release links that are required for remote release of the mooring and retrieval from the sea
288 surface. Many of the releases (located at 10 m above seafloor) used in the Monterey CCE did
289 not release properly when the command was issued from the support vessel. The extreme case
290 was the final deployment where every mooring required a Remotely Operated Vehicle (ROV)
291 dive to recover the mooring. Some of these required cutting of the mooring string below the
292 release, while others only required tapping the release with the ROV's mechanical arms. These
293 issues were attributed to the presence of sand within the releases and are similar to those
294 encountered by single-point moorings in the submarine channel in Bute Inlet, where Prior et al.
295 (1987) recorded: i) damage, removal and fouling of rotors and vanes on current meters (causing
296 poor data quality); ii) bent and sheared shackles and stainless steel frames; iii) up to 1 km down-
297 channel transport of moorings; iv) failure of acoustic releases to detach due to burial by sand;
298 v) parting of mooring lines; and vi) the entire loss of some instruments (also detailed in Zeng et
299 al., 1991).

300 Unlike these examples from Monterey Canyon and Bute Inlet, no irreversible pressure or
301 temperature changes were observed for the single-point mooring in the Congo Canyon.
302 Therefore the Congo Canyon mooring is unlikely to have been moved by any of the eleven
303 turbidity currents that occurred during its deployment (Azpiroz-Zabala et al., 2017).
304 Furthermore, no damage was recorded in this case to either the acoustic release links or the
305 ADCP. No other instruments were placed on the mooring line.

306



307

308 **Figure 3: Photographs of sediment trap and in-line instruments placed within turbidity**
 309 **currents from Monterey Canyon. A: Pre-deployment photograph of sediment trap and**
 310 **instruments fitted on cantilevered aluminium brackets. B: Deployment of sediment trap.**

311 **C: Detail on anchor weight (train wheels) and acoustic release links, which were placed 3**
312 **m above seafloor. D & E: Sediment trap and instrument brackets following retrieval,**
313 **showing damage and fouling during interaction with turbidity currents.**

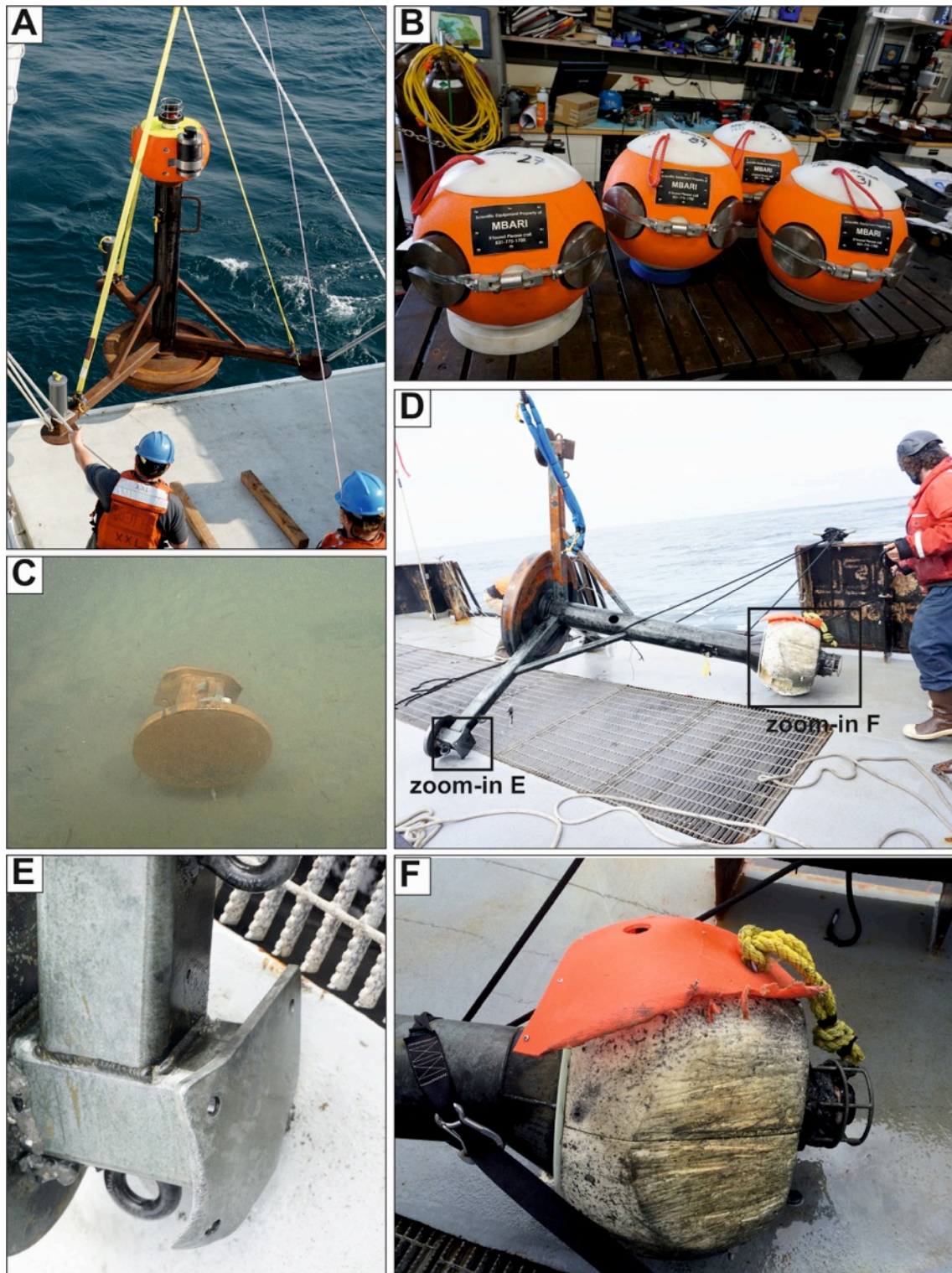
314

315 **3.3. Burial, down-slope transport and damage of seabed frames**

316 We now discuss issues that have affected seabed-based platforms. An upward-looking ADCP
317 was mounted on a bottom-mounted tripod in 2011 and deployed at the terminal end of a
318 submarine channel offshore from the Squamish river delta (150 m water depth). This ADCP
319 recorded 22 turbidity currents of up to ~ 1.5 m/s over a period of four months (Hughes Clarke et
320 al., 2012), with the exception of a 20 day period when the run-out from a delta-lip collapse led
321 to the burial of the frame (Clare et al., 2016). With a single ensemble averaging interval of 20
322 seconds, the ADCP went from recording flow to being completely buried. Thus, no monitoring
323 was possible during this time. Interestingly the ADCP frame was not significantly tilted in this
324 process. Fortunately a vertically offset surface buoy was attached so that the instrument could be
325 dragged out of the sediment.

326 In addition to the movement of single-point moorings deployed in the Monterey CCE, down-
327 canyon movement of an 800 kg AMT-tripod-frame (Figure 4A) was also recorded six-times.
328 These episodes of movement corresponded to the timing of turbidity currents. On the 15th
329 January 2016, the AMT frame moved 4.2 km down-canyon and was observed from ROV video
330 to be on its side, half-embedded within in the seafloor (Paull et al., 2018). Following its
331 redeployment, the mooring was transported 0.9 km on 24th November 2016; also found on its
332 side, but this time buried by at least 2 m of sediment with only one foot protruding at seafloor
333 (Figure 4C). The heavy-duty steel frame was sand-blasted, its feet bent and sheared in places,
334 while much of the pressure-resistant foam coating was abraded from the Benthic Event Detector
335 (Figure 4D-F). Pressure, temperature and accelerometer measurements indicate that once the
336 AMT frame was tilted onto its side it became buried during the initial turbidity current, and then
337 remained in that position, until it was moved by successive flows. A multi-instrument 'Seafloor
338 Instrument Node' (SIN) was placed in a deeper water location (1840 m), where the Monterey
339 Canyon widens. Flows decelerate from ~ 4 -8 m/s in the upper part of the canyon where the AMT
340 frame was deployed to ~ 1 -2 m/s at the SIN location (Figure 5; Paull et al., 2018; Heerema et al.,
341 2019). Impacts of turbidity currents were less severe at this more distal location; however, the
342 SIN frame was also transported down-canyon, by 26 m, and now was locally buried by up to 34
343 cm of sediment (Figure 5C). A high frequency acoustic instrument (Aquadopp) was ripped from

344 the arm that suspended it above seafloor and up to 10 cm of scour was noted from repeated
345 ROV-based bathymetric surveys (Figure 5B&C).

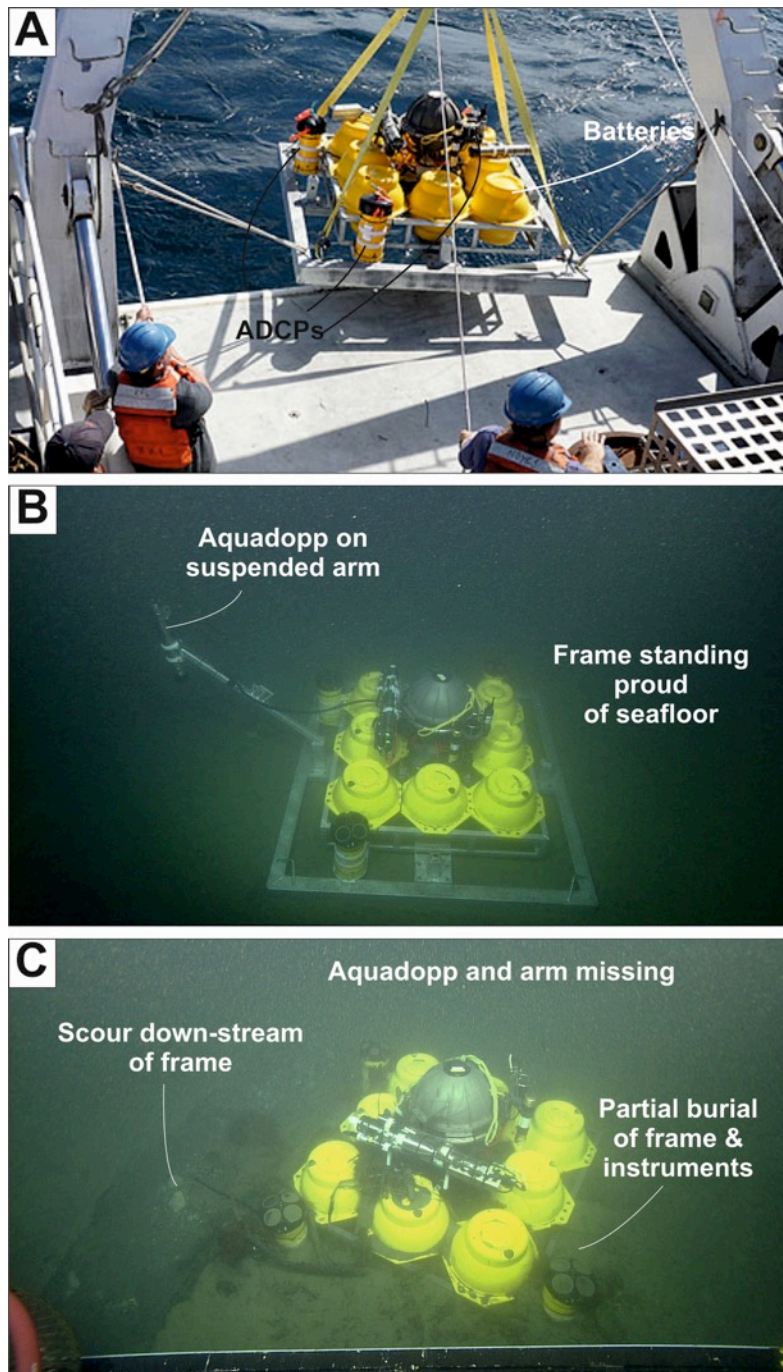


346

347 **Figure 4: Photographs of the 800 kg AMT frame deployed at 300 m in Monterey Canyon.**

348 **A: Prior to deployment of instrument. B: Example of Benthic Event Detectors, one of**

349 which was attached to the top of the AMT frame to track the sense of motion of the frame.
350 C: Only the foot of the AMT frame was found protruding from seafloor by ROV dive
351 video following its burial by a turbidity current. D: AMT frame following retrieval to
352 deck, revealing damage to the frame (E) and the Benthic Event Detector (F) caused during
353 its down-canyon transport.



354

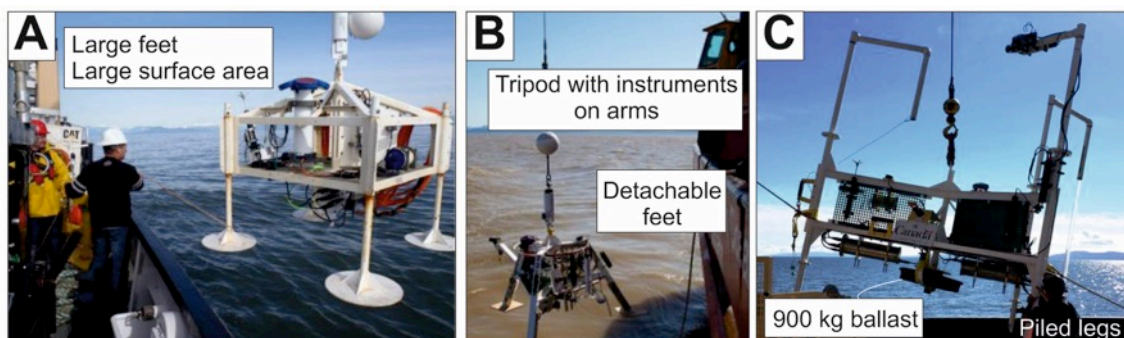
355 **Figure 5: Photographs of the Seafloor Instrument Node (SIN), deployed at 1840 m water**
356 **depth in Monterey Canyon. A: SIN prior to deployment. B: ROV video still showing**

357 **deployed location where the frame sits proud of seafloor. C: ROV video still at retrieval,**
358 **following 26 m down-canyon transport, with evidence of local scour and deposition around**
359 **the frame and removal of the Aquadopp and its mounting arm.**

360

361 Even benthic landers sited outside of submarine channels can suffer from adverse impacts that
362 include burial and movement of the platform. The original Delta Dynamics Laboratory platform
363 (DDL), deployed in 2008 on the Fraser Delta (located in a bedform field outside of the main
364 submarine channel; Figure 2E), was buried by as much as 1 m of sediment. Initially it was
365 thought that this was simply natural sediment deposition from the Fraser River; however it is
366 now attributed to active turbidity currents (Lintern et al., 2016). Recovery using a vessel-
367 deployed crane caused a large ship (the 1800 tonne CGS John P. Tully) to lean uncomfortably
368 and snapped 9,000 kg lines. The original platform design at the Fraser Delta had a large surface
369 area, which also made it prone to tumbling during turbidity currents, as recorded by frame-
370 mounted orientation sensors, and was therefore replaced by a lower-profile platform with
371 weighted legs (Figure 6; Lintern et al., 2016). This revised deployment included arms and poles
372 that held instruments away from the platform and above the 2 m powerful flows that were
373 detected (Lintern et al., 2019). It also featured feet that snap free on retrieval, as embedment of
374 the original large feet created problems during recovery (Figure 6B). The second platform
375 design mostly remained upright, but sometimes slid downslope during strong turbidity currents.
376 To make it more resistant to flows, over 900 kg of ballast is suspended below the platform,
377 while the legs penetrate the seafloor by up to 1 m, acting as small piled foundations (Figure 6C).
378 This enhanced design has so far remained upright for two years, experiencing flows of up to 9
379 m/s (Lintern et al., 2017; 2019).

380



381

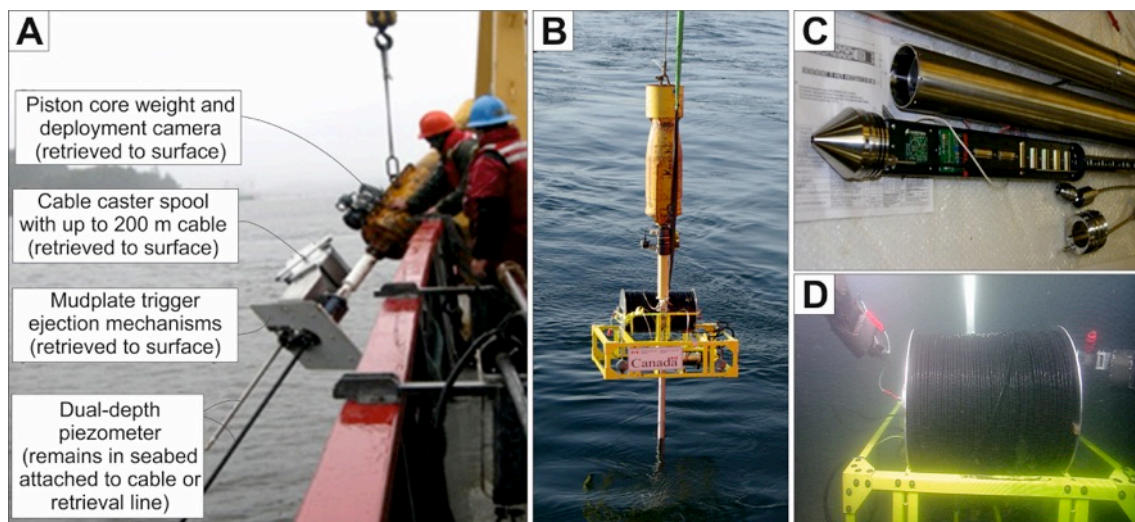
382 **Figure 6: Development of the Fraser Delta Dynamics Laboratory including A)**
383 **conventional design with large feet to stop embedment, B) revised tripod design with**
384 **detachable feet. Both A and B tumbled down-slope during powerful flows. C) Revised**

385 design that has withstood numerous powerful flows to date due to its piled legs and
386 ejectable ballast weight. Image modified from Lintern et al. (2019).

387

388 Other platforms on the Fraser Delta include a benthic boundary laboratory (BBL) and a seismic
389 liquefaction in situ penetrometer (SLIP; Figure 7). The BBL's main design feature was a
390 cantilever to hold instruments away from the main platform to minimize frame turbulence.
391 Despite the increased tipping moment this would appear to cause, it is worth noting that further
392 down the delta slope (140 m versus the DDL 107 m) there has not been a strong enough
393 turbidity current in five years of deployment to topple the deeper BBL. The SLIP is an
394 instrument designed to measure pore pressures which could be associated with subaqueous mass
395 movements (Figure 7). It is constructed of a fibreglass frame above the seafloor holding a
396 system of valves, data-loggers, instruments, and a network plug. The data logging is done on
397 cyclical buffers and has backup battery power in case of being severed from the network. The
398 lower part of the SLIP is a 5 m-long cone tip with multiple pressure and temperature ports. An
399 800 kg piston core head weight is used to push the SLIP tip into the sediment. The SLIP has
400 been deployed for several years at the site of the DDL, and due to its 5 m embedded tip, it has
401 not suffered any translation from the same turbidity currents, which have been tumbling the
402 DDL platforms.

403



404

405 **Figure 7: The prototype Seismic Liquefaction In Situ Piezometer (SLIP) at the Fraser**
406 **Delta. A: Overview of instrument prior to deployment. The large stainless steel container**
407 **houses data processing and logging instruments, and an underwater modem. All**
408 **components are made from fibreglass or stainless steel in an attempt to minimize**

409 corrosion in salt water. Power is provided by the network, and data is transmitted directly
410 to the scientists' offices over the internet in near real time. A battery backup and circular
411 buffer continue to measure data in case of a severed cable, due to a slope failure. B:
412 Deployment using 816 kg weight. C: Detail on instrumented tip that contains devices to
413 measure earthquakes and ground movements, measuring up to 100 times per second. D:
414 The cable being unspooled 1.5 km to the Victoria Experimental Network Under the Sea
415 (VENUS) node by the manipulator arms of an ROV.

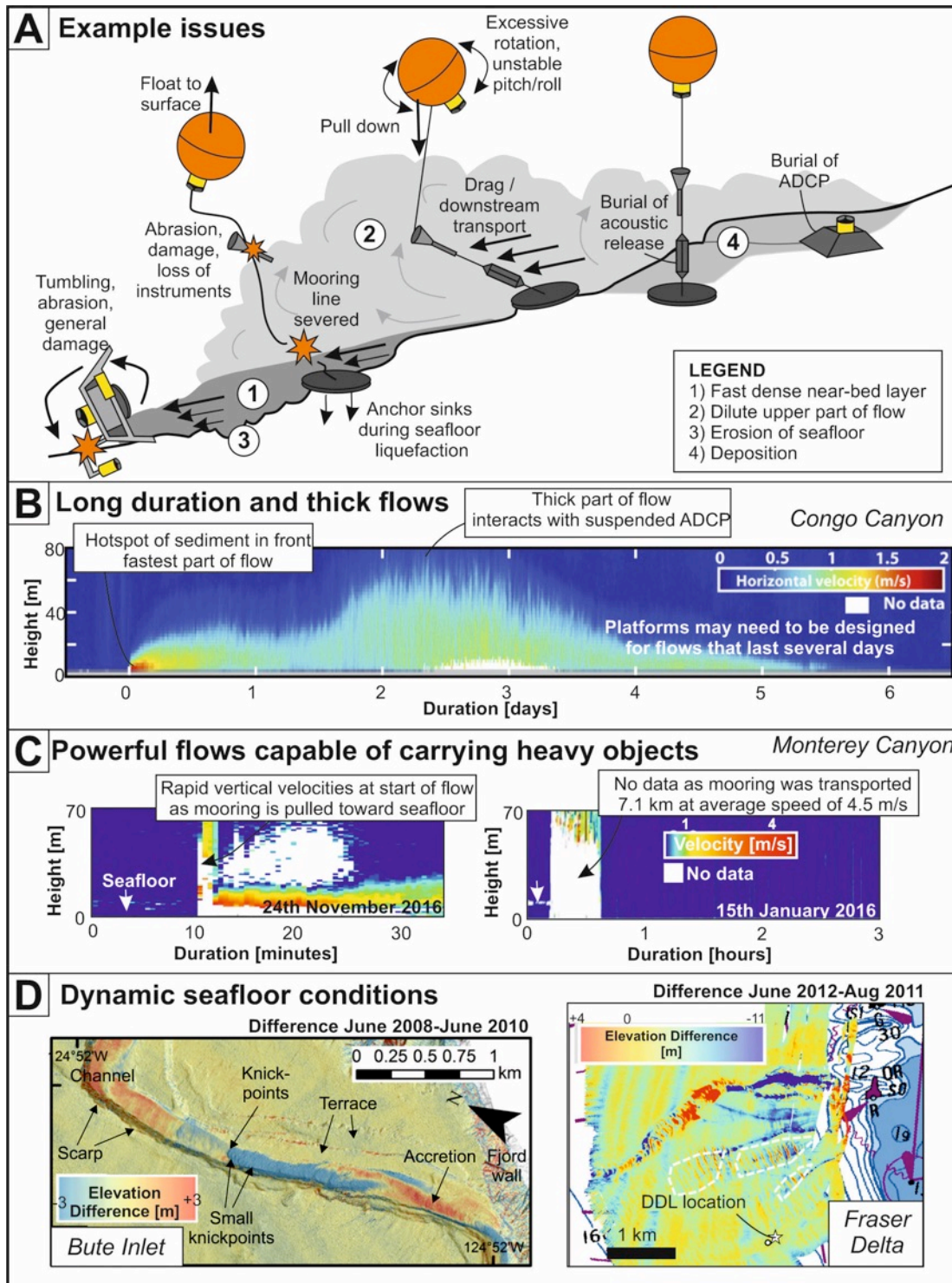
416

417 3.4. Overview of adverse impacts related to turbidity currents

418 Based on past experiences from recent monitoring campaigns, the following observations can be
419 summarised about the hazards posed by turbidity currents to moorings and seafloor platforms
420 (Figure 8):

- 421 1) The powerful dense near-bed part of a turbidity current (particularly prone in proximal
422 confined submarine canyons or channels) may be capable of toppling and/or
423 transporting heavy (>100s of kg) objects, including anchors and seabed frames (Figure
424 8A). This dense part of the flow can damage platforms, sensors and ancillary mounting
425 equipment through collisional impact or drag, and may even result in short-lived
426 liquefaction of seafloor sediments, causing anchors for single-point moorings or
427 seafloor frames to sink.
- 428 2) Fast flows may pull instruments down towards seafloor, and in some cases overcome
429 the tractional forces required to keep the anchor in place, and transport single-point
430 moorings down-channel (Figure 8A).
- 431 3) Where instruments interact with a turbidity current, this may lead to platform instability
432 and poor quality data, damage to acoustic releases (jeopardising successful retrieval of
433 moorings) or, in severe cases, loss of instruments and mooring components (Figure 8C).
- 434 4) Erosion of the seafloor may change local seafloor elevation and undermine platforms
435 where it occurs as scour around a seafloor structure (Figure 8D&E).
- 436 5) Sudden deposition, sometimes involving several metres thickness of sediment, can bury
437 seafloor instruments or low-elevation acoustic releases, limiting instrument
438 performance and causing issues for retrieval (Figure 8D&E).

439



440

441 **Figure 8: (A) Overview of some of the issues encountered in monitoring active turbidity**
 442 **currents discussed in this paper. (B) An example of one of the long-duration turbidity**
 443 **currents measured in the deep-water Congo Canyon that may attain thicknesses of >80 m**
 444 **(modified from Azpiroz-Zabala et al., 2017a). (C) Two turbidity current events measured**
 445 **at the shallowest water mooring in the Monterey Coordinated Canyon Experiment in**

446 **Monterey Canyon. On the left is a flow that pulled the instruments and buoyancy towards**
447 **seafloor at the start of the event due to enhanced drag early on. On the right is the record**
448 **from an ADCP that was transported by a flow at several m/s; hence no reliable data were**
449 **recorded during the flow. This mooring was transported 7.1 km down-canyon and then**
450 **broke free from its anchor and was released to the sea surface. (D) Repeat multibeam**
451 **echo-sounder seafloor surveys illustrating how active turbidity currents can both erode**
452 **and deposit at seafloor. The location of the Delta Dynamics Laboratory is labelled on the**
453 **Fraser Delta (right).**

454

455 **4. Designing monitoring platforms to successfully measure turbidity currents**

456 In this section we highlight some of the lessons we have learned from previous turbidity current
457 monitoring campaigns, to inform future ones.

458

459 **4.1. Finding a ‘sweet spot’ for the design of single-point moorings**

460 When optimising mooring design to address one issue, other complications may arise
461 concerning another. We now discuss how mooring designs have been iteratively refined to try
462 and find the ideal configuration for different settings and objectives.

463

464 **4.1.1. Reduce the surface area to minimise drag**

465 Single-point moorings are typically the preferred way to monitor turbidity currents as they can
466 be deployed from the back deck of an ocean-going vessel equipped with a suitable winch and A-
467 Frame. Successful monitoring of turbidity currents is strongly dependent on the mooring design.
468 Single-point moorings in the Congo Canyon did not show any movement down-canyon during
469 turbidity currents, nor were any of the instruments damaged. We identify three reasons for the
470 stability of this Congo Canyon mooring. First, while the flows in Congo Canyon lasted many
471 hours to days in duration, they were generally muddy and dilute flows (with the exception of a
472 frontal cell of sand-rich sediment-laden fluid), and relatively slow, reaching maximum
473 velocities of <3 m/s with an average of <1 m/s (Azpiroz-Zabala et al., 2017a). Conversely,
474 flows in the Monterey Canyon, often reached velocities far in excess of this value; up to 7.2 m/s
475 and are interpreted to have been denser, with the near-bed part of the flow capable of
476 transporting gravel and cobble-sized material (Paull et al., 2018). Second, the mooring design
477 for the Congo Canyon included heavier anchor weighting (~2000 kg), use of low-drag neutrally

478 buoyant plastic-coated mooring line and a larger syntactic buoy housing the ADCP. This greater
479 buoyancy ensured the mooring line remained taut during flows (Figure 9). Third, and perhaps
480 most importantly, the mooring design was much simpler for the Congo Canyon measurements
481 than in Monterey Canyon (Figure 9). Sediment traps were not deployed, and acoustic release
482 links were placed far (~40-60 m) above the velocity maximum of the flows, in order to reduce
483 drag on the mooring line imparted by flows. Therefore, one way to maximise the likelihood of
484 successful monitoring is to ensure that any instruments are located above the turbidity currents
485 that you wish to observe, which will decrease the likelihood of drag and also add weight to the
486 mooring line. Previous successful deployments in the Var Canyon (Mediterranean) used lower
487 frequency (75 kHz) downward-looking ADCPs that were placed much higher (300-350 m)
488 above seafloor than the higher frequency 300-600 kHz instruments in the Monterey and Congo
489 Canyons (Khrifounoff et al., 2012). Coarser vertical resolution was accepted to ensure that the
490 single-point moorings interacted less with turbidity currents. The Var Canyon deployments also
491 featured ADCPs set within gimballed frames that ensure the ADCP can tilt to remain as vertical
492 as possible. Such a situation may be unavoidable, however, if you wish to: i) measure close to
493 the seafloor using high frequency instruments (e.g. Hughes Clarke et al., 2012; Clare et al.,
494 2015; Hughes Clarke, 2016); ii) sample sediments within the flow to measure vertical grain size
495 segregation or quantify organic particulate flux (e.g. Maier et al., 2019a&b); iii) make
496 measurements within the flow to ground-truth other remote sensing style measurements (e.g.
497 Azpiroz-Zabala et al., 2017a; Hage et al., 2019). Sediment traps are typically the largest item on
498 the mooring line; hence its height above the bed may be critical. The style of sediment trap also
499 makes a difference. Mclane-type traps provide a greater cross-sectional area than the narrower
500 Anderson-type traps.

501

502 **4.1.2. Design anchor weight and flotation appropriately, particularly if multiple** 503 **instruments are required for single-point moorings**

504 One of the primary goals of the Monterey Canyon CCE was to estimate suspended sediment
505 concentrations during a turbidity current using the acoustic backscatter from the downward
506 facing ADCP. Given that the acoustic response of the ADCP is both a function of the
507 concentration and the grainsize of the material in suspension, it was decided that an in-line
508 sediment trap was essential, even if the presence of the trap increased drag on the mooring.
509 While it may seem intuitive that increasing the anchor weight will improve mooring stability,
510 this is not always the case. Moorings deployed in Monterey Canyon in the early 2000s had
511 multiple train wheels for their anchor and long mooring lines with multiple instruments attached

512 (Xu, 2011). Some of these moorings were lost due to the drag exerted during turbidity currents
513 and the mooring line parted. Conversely, a mooring has been deployed successfully at 1300 m
514 water depth in the Monterey Canyon, almost continuously since 2002 with minimal ballast
515 (scrap steel) (Barry et al., 2006; Xu et al., 2013). Its light ballast makes this mooring relatively
516 easy to move, but this also ensures that the strain on the mooring line does not reach a critical
517 limit. Thus, one way for a mooring to survive may be to allow it to be dragged down canyon.
518 This philosophy is also in keeping with minimising the amount of debris that is left behind
519 following mooring retrieval, as it is difficult to justify leaving iron, cables and potentially
520 fibreglass in the marine environment.

521

522 There appears to be a 'sweet spot' for mooring design that involves a compromise between
523 minimising drag (which may not be possible if several instruments need to be deployed within
524 the flow height), stabilising the mooring base with anchor weight, and maximising buoyancy to
525 vertically stabilise the mooring line. The design of the mooring is an iterative process, balancing
526 available anchor weight, surface drag (and weight) from in-line mooring elements, and both in-
527 line and top flotation elements. The mooring design toolbox written in Matlab by Richard
528 Dewey (Mooring Design & Dynamics; Dewey, 1999) was used in the Monterey Canyon CCE to
529 evaluate the performance of the single-point taut-wire moorings. The program allows a user to
530 design a surface or subsurface wire mooring, and contains a large database of the physical
531 characteristics of standard oceanographic equipment (such as dimensions, submerged weight,
532 surface drag), and will evaluate how a mooring responds to a static flow profile (i.e. does not
533 account for waves). It was thought that turbidity currents in Monterey Canyon did not exceed 2
534 m/s (since the most recent data derived from one hour averages), and this value was used in the
535 initial mooring designs. We now know this was a considerable underestimate. It is best to have
536 contingency and overdesign. Keeping the mooring as upright as possible (increasing the in-line
537 tension) required additional flotation (Figure 12A), which has the additional negative effect of
538 making the anchor 'lighter' by increasing the upwards force on the anchor, thereby making the
539 mooring more likely to move down-canyon during events. Even the type of flotation used was a
540 consideration: in shallower water (less than 800 m) plastic flotation was used for in-line
541 elements, to provide greater flotation per diameter of sphere (and thus surface area, because they
542 weigh less) than comparably sized glass or syntactic foam elements. Increasing the anchor
543 weight from 450 kg to 900 kg, and increasing the in-line flotation above each sediment trap and
544 the top of the mooring, dramatically improved mooring performance demonstrating that it is
545 possible to refine the design successfully. This may require some *a priori* knowledge of the
546 likely flow conditions. Regardless of design, one key lesson learned is to include an iridium

547 beacon on the instrument package such that it can be tracked should it cut loose and float to the
548 surface.

549

550 **4.1.3. Strengthen the weak points on a monitoring platform: strategic placement of**
551 **acoustic releases and resilient instrument mounting**

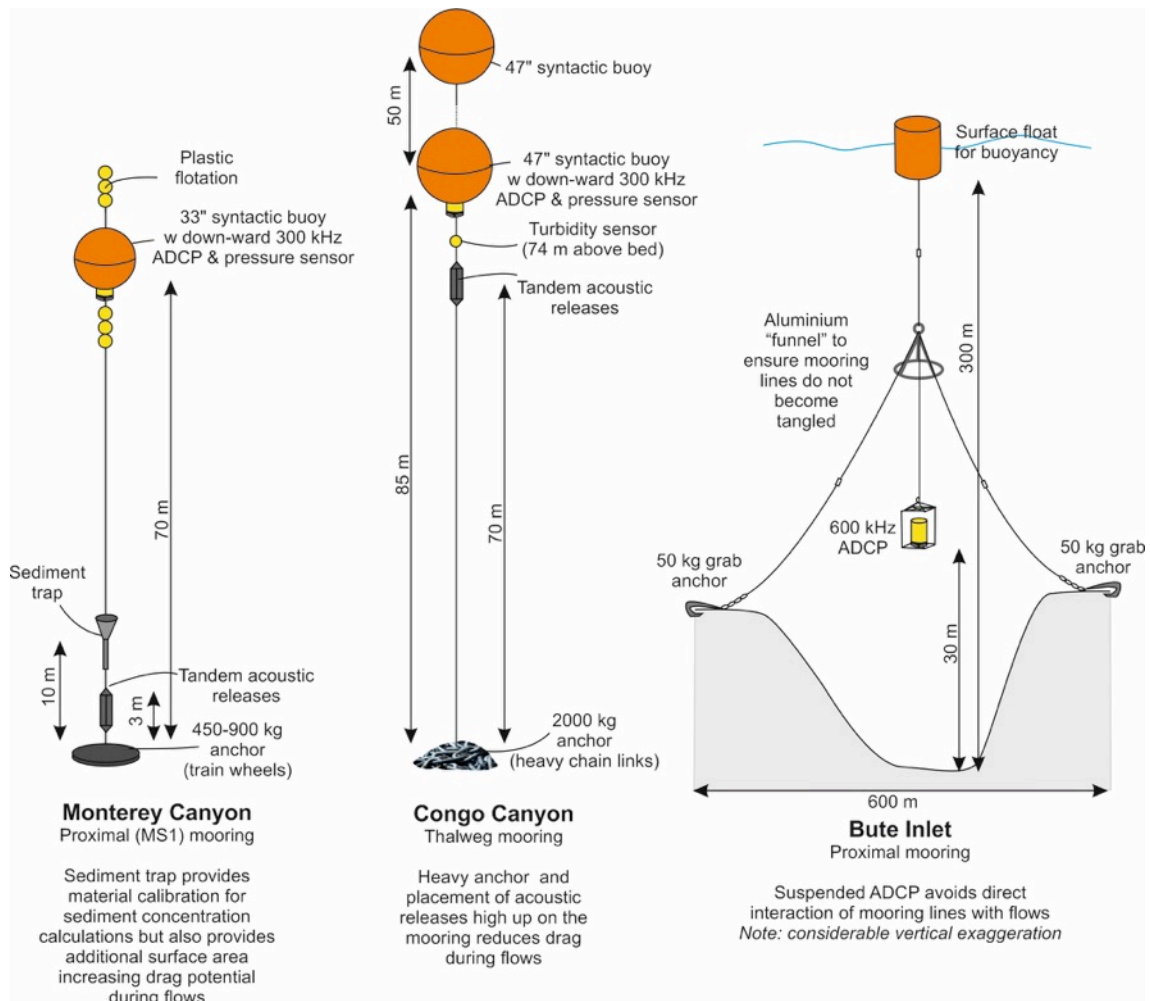
552 Had the Monterey Canyon CCE not been supported by an ROV, then the failure of the acoustic
553 releases (placed close, 3 m, above the seafloor) to return the moorings to the surface would have
554 meant the loss of valuable data and instruments. Many research and industrial expeditions do
555 not have the benefit of a support ROV; hence, we recommend that acoustic release links are
556 placed as high as practicable above seafloor, where they are away from the damage that may be
557 caused at the sand-rich base of a turbidity current (but low enough such that they do not
558 interfere with the ADCP). A recent study in the Gulf of St Lawrence (E Canada) by
559 Normandeau et al. (2019a) suggested placing the acoustic release a minimum of 1 m above the
560 height of intra-channel bedforms, to avoid interaction with the most vigorous and potentially
561 dense part of the flow. Tandem acoustic release links are routinely deployed for single-point
562 moorings (i.e. to provide redundancy in case one fails) but it may also be sensible to deploy the
563 releases in series, rather than in a parallel twinned deployment so that they are not both subject
564 to impacts at the same elevation within the flow (Xu, 2004).

565

566 Instrument mountings were often found to be weak points in a monitoring platform's design
567 (e.g. Figure 3&4). In the case of the Monterey CCE deployments, near-bottom current meters
568 and altimeters (10 m above seafloor), were mounted on protruding brackets (cantilevered) on
569 the single-point moorings 1 m from the sediment trap strong-back with ¼" aluminium angle
570 stock (instead of stainless steel, to reduce weight; Figure 3). It was necessary to cantilever them
571 away from the mooring in order to ensure that instruments below the ADCP were not affected
572 by the mooring wire, or other instruments below. This design provided an even larger surface
573 area for drag and also increased the weight on the mooring line, however, and underlines how
574 operational necessities may end up going against the guidance to minimise drag. The aluminium
575 design survived four turbidity currents, but eventually broke. In future, and if resources allow,
576 we suggest that titanium should be used for mounting in similar environments. Heavy metal
577 parts and coated iron wires should be avoided, especially for long-term deployments, as it is
578 impossible to have a visual check on corrosion. Instead, plastic-coated Ultra High Molecular
579 weight polyethylene Dyneema rope is preferred as there are no corrosion issues, they are thin
580 and neutrally buoyant, and may be used for multiple deployments.

581 While it may be possible to strengthen brackets and frames, any instruments with moving
 582 external parts (e.g. the impellers that were damaged on the current meter deployed in Monterey
 583 Canyon) or that protrude away from the platform (e.g. the steel arm that held the near-bed
 584 Aquadopp in Monterey Canyon; Figure 4) are likely to be vulnerable and should be considered
 585 to be at high risk during field deployments to measure powerful flows.

586
 587



588

589 **Figure 9: Comparison of subsurface single-point moorings deployed in Monterey Canyon**
 590 **(Paull et al., 2018), Congo Canyon (Cooper et al., 2012; Azpiroz-Zabala et al., 2017) and**
 591 **two-point mooring supported by surface buoy in Bute Inlet. Not drawn fully to scale.**

592

593 **4.2. Suspended monitoring systems that avoid instrument and mooring-line interaction**
 594 **with the flow**

595 To avoid the damaging effect of a passing turbidity current (e.g. drag, scour, burial), another
596 option is to avoid placing instruments, anchors and mooring lines within the flow at all. Such an
597 approach may also be necessary where the available support vessel for deployment cannot
598 handle the bulky hardware (e.g. c.1 m diameter syntactic buoys and stack of train wheels
599 weighting c.1 tonne) required for single-point moorings. We now discuss two plausible
600 geometries: i) hull mounted systems; and ii) surface buoy suspended systems with two or more
601 anchors. Both of these are only practical in shallow water (typically <500 m) environments,
602 given the amount of deck space used and the logistics involved with such quantities of mooring
603 line and anchors, and for short-term (months) deployments. Such methods are therefore only
604 generally applicable in fjord or lake environments, and not the deep ocean; however previous
605 deployments in the Var Canyon has demonstrated that subsurface two-point moorings are
606 feasible in water depths as great as 1280 m (Khripounoff et al., 2012).

607

608 **4.2.1. Vessel-mounted monitoring systems**

609 *Hull-mounted deployments*

610 Hull-mounted systems include acoustic imaging (downward looking single or multibeam sonars
611 or ADCP) and rapidly descending underway physical probes (e.g. Moving Vessel Profiler,
612 MVP; Hughes Clarke et al., 1996). For any of the sonar systems, the issue becomes resolution –
613 the further away from the seafloor, the poorer the range resolution usually is (longer, narrow-
614 band pulses required); especially the angular resolution. For single beam sonars the width of the
615 projected beam (typically 7-30 degrees) may result in echoes from offset roughness elements
616 (like bedform crests or channel flanks) which can be confused with the real near-seafloor
617 profile. Multibeam systems (with beam widths in the 1-2 degree range) provide far better
618 definition (See Figure 10A-C; Hatcher, 2017). For ADCPs, just as with the conventional
619 downward-looking single-point moorings, the closest usable data to the seabed is limited by the
620 first echo of the projected side lobes from the beams inclined at 20 degrees (Figure 12B). This
621 limits the first usable bin to about 10% of the ADCP altitude (using conventional 4-beam
622 systems). Thus, to investigate 5 m thick flows for example, surface-mounted ADCPs would not
623 be of use at elevations much greater than ~ 50 m, plus the vessel has to be present at the time of
624 the flow.

625 Therefore, this surface-mounted method is only viable if the flows are known to be frequent
626 and/or of known likely timing. This was the case for the Squamish 2011-2013 and 2015
627 campaigns (Hughes Clarke, 2016). Here, a small vessel (CSL Heron) deployed an MVP. The
628 MVP consisted of a tow body with a conductivity, temperature and depth (CTD) and an optical

629 backscatter probe that can be released at slow speeds (< 6 knots). If the vessel slowed down for
630 the descent duration (typically 2 minutes) the probe descended to a depth of 100 m. The MVP
631 was deployed daily along the main channel sections to catch evidence of suspended sediment
632 clouds due to a passing turbidity current. On a few occasions, the MVP was able to sample the
633 top of an active turbidity current, which was also observed in the EM710 water column imagery
634 (1x2 degree beam, 0.2 to 0.5 ms pulses, 70-100 kHz; Hughes Clarke et al., 2014; Hage et al.,
635 2019).

636

637 The MVP has several limitations. The profile is necessarily discrete. The minimum horizontal
638 spacing depends on the time it takes to winch back in the instrument cable, typically 5 minutes
639 if going to 100 m. The instrument package is deliberately designed to stop free-falling before
640 hitting the seabed. Thus measurements closer than 5 m from the actual seabed are rare, and only
641 the top of an active flow is usually recorded. The use of hull-mounted instrumentation will only
642 be useful in relatively shallow water where the recurrence of active turbidity currents is
643 reasonably predictable. This is not the case for most turbidity current systems, where longer-
644 term un-crewed campaigns are required.

645

646 *AUV-mounted deployments*

647 Autonomous Underwater Vehicles (AUVs) now enable the acquisition of high-resolution
648 seafloor datasets, by flying the AUV close to seafloor (Wynn et al., 2014). These autonomous
649 mobile systems can also hold instruments, such as ADCPs, to monitor the seafloor along
650 transects, in the same manner as river systems are often measured (e.g. Parsons et al., 2007). A
651 saline density underflow has been monitored using such an approach, to the north of the
652 Bosphorus Strait in the Black Sea. Along- and across-channel transects of ADCP measurements
653 were acquired using a 1200 kHz ADCP, revealing a range of flow dynamics, which include
654 evidence for secondary circulation cells and the presence of hydraulic jumps. These jumps had
655 previously only been hypothesised from laboratory experiments of submarine channels (Parsons
656 et al., 2010; Sumner et al., 2013; Wynn et al., 2014; Dorrell et al., 2016; Azpiroz-Zabala et al.,
657 2017b). This AUV-based monitoring was also performed in a very narrow and busy shipping
658 lane; hence surface-based monitoring would have been precluded (Wynn et al., 2014). Future
659 developments in AUV endurance (e.g. battery performance) may make this type of monitoring
660 more common, however, it is only likely to be used where the timing of the flow is very well
661 constrained or continuous, as in the case of the Bosphorus underflow.

662

663 **4.2.2. Multi-point anchoring for vessel-based monitoring**

664 If a turbidity current is laterally restricted by canyon or channel flanks, it is possible to use two
665 or more anchors located on either side of the channel to position a surface buoy above the active
666 channel, from which a variety of instruments can then be suspended. In practice, there are depth
667 limitations to this, as the longer the anchor lines, the more the suspended instrument is likely to
668 move. The first test of the two point anchoring method, occurred in 2014 at 200 m water depth
669 in Bute Inlet, and then in 2015, in Squamish in 120 m of water. From 2017 to 2019, a two-point
670 mooring was deployed in Squamish at the lobe channel termination in 160 m of water (Figure
671 10D). A minimum of two anchors can adequately constrain the buoy across the channel, but any
672 slack in the lines will allow the buoy to move slightly along the channel as a result of wind or
673 tide drag on the surface buoy and the suspended lines. A third anchor helps constrain the along
674 channel motion.

675

676 Four-point moorings were deployed for the 2013 Squamish experiment (Hughes Clarke, 2016).
677 This four-anchor approach not only best constrained the surface location, but also allowed the
678 suspended instruments to be held at a fixed azimuth. For any number of anchors, if there is only
679 a single surface buoy, the suspended instrument is free to rotate in azimuth. Thus the instrument
680 measurement must not be compromised by this rotation. Instruments which have an internal
681 compass can correct for such rotations; however, any system that requires at a preferred azimuth
682 (such as the acoustic monitoring of a fixed stretch of channel by a forward-looking multibeam
683 (M3) imaging used in the 2013 and 2015 Squamish experiments; Hughes Clarke, 2016; Hage et
684 al., 2018) would not be usable. To overcome this problem, in the 2013 Squamish experiment,
685 the four anchors were arranged in pairs to come up to two surface buoys located offset along the
686 channel below. The surface buoys were in turn held together by a surface line. The vessel tied
687 up daily between the two buoys and azimuth sensitive instruments (the M3 sonars described in
688 Hughes Clarke, 2016), were suspended on a frame that was attached fore aft so that it could not
689 rotate significantly in azimuth. Such a deployment is only suited to short-term (days to weeks)
690 duration.

691

692 **4.2.3. Two-point moorings for autonomous deployments**

693 In 2016, and again in 2018, two-point moorings were deployed in water depths of up to 450 m
694 in Bute Inlet (Figure 9). Such a mooring design was conceived to remove any drag on the

695 instruments or mooring line and because previous repeat seafloor surveys indicated that
696 turbidity currents may be capable of depositing and eroding up to tens of metres of sediment
697 (Conway et al., 2012; Gales et al., 2019). In this example, the anchor lines no longer
698 connected to the surface float, but instead to a triangular frame that was suspended at about half
699 the water depth from the surface float (Fig. 9, 11C&D). This setup reduces the length of the
700 anchor lines, limiting the movement of the instrument and facilitates the deployment. With the
701 submerged frame at half the water depth, the anchors can be placed one at a time and the anchor
702 can be dropped with slight tension in the line at the final anchor location. Each anchor and chain
703 had a weight of 100 kg, while the surface float carrying the instruments had a buoyancy of 150
704 kg. The anchor lines were about 300 m to keep the anchors well away from the 200 m wide
705 channel (Figure 9). Such a deployment is logistically challenging, particularly on smaller
706 vessels, and is therefore unlikely to become a routine mooring configuration. The use of a
707 surface buoy would also be impractical in high latitude settings where seasonal sea ice forms. A
708 completely submerged two-point mooring was deployed in the deep-sea Var Canyon, however,
709 which involved anchoring either side of the canyon in a water depth of 1200 m (Khripounoff et
710 al., 2012). Acoustic releases were placed 10 m above the anchor on each of the mooring lines,
711 so that only 10 m lengths of rope and the anchors were left on the seafloor following retrieval.
712 Large quantities of mooring lines, buoyancy and anchors meant that this mooring required a
713 spacious back deck of a large ocean-going vessel.

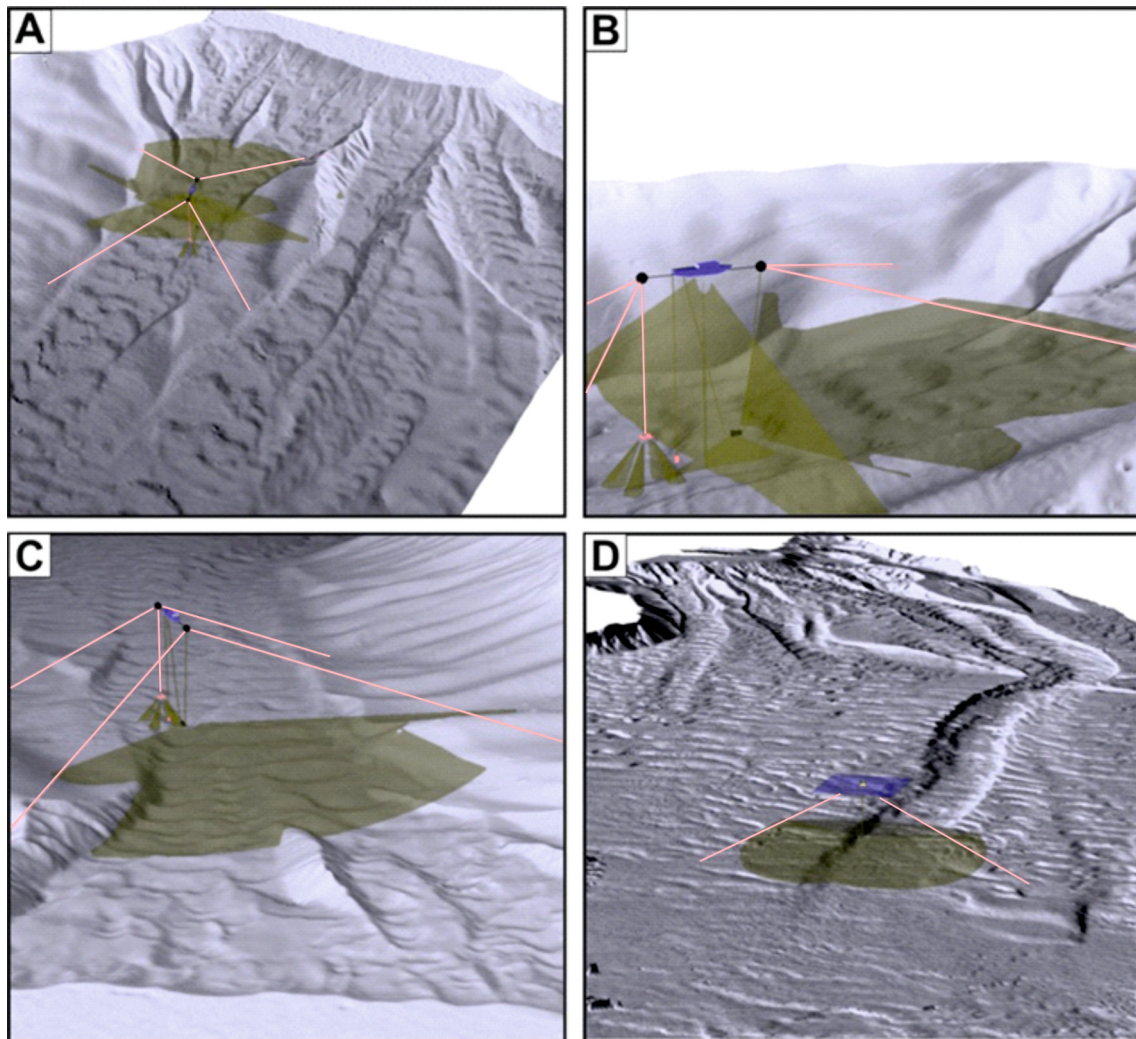
714

715 **4.2.4. Decision on instrument elevation for multi-point moorings**

716 A two-point mooring with a surface buoy will only fix the elevation of the ADCP with respect
717 to the surface. Any tidal excursions will result in the ADCP moving toward and away from the
718 seabed. Deciding on the optimal height may require prior information on the likely thickness of
719 the flows. The instrument should be above the active flow, even if partly in the turbid cloud of
720 the wake, but at the same time, as close to the seabed as possible so that side lobe masking is
721 minimized (Figure 12B). In larger systems and deeper canyons, where the anchors for a two-
722 point mooring are designed to be above the height of the flow (e.g. due to their location on
723 canyon terraces), the ADCP will be at a considerable height above the seafloor; out of the range
724 of high frequency instruments such as 300 or 600 kHz ADCPs. A two-point mooring
725 configuration in the Var Canyon enabled the first monitoring of powerful turbidity currents and
726 a debris flow, with thicknesses sometimes in excess of 100 m (Khripounoff et al., 2012). At this
727 height a 75 kHz ADCP, placed >300 m above seafloor, was necessary to have sufficient range to
728 capture this flow.

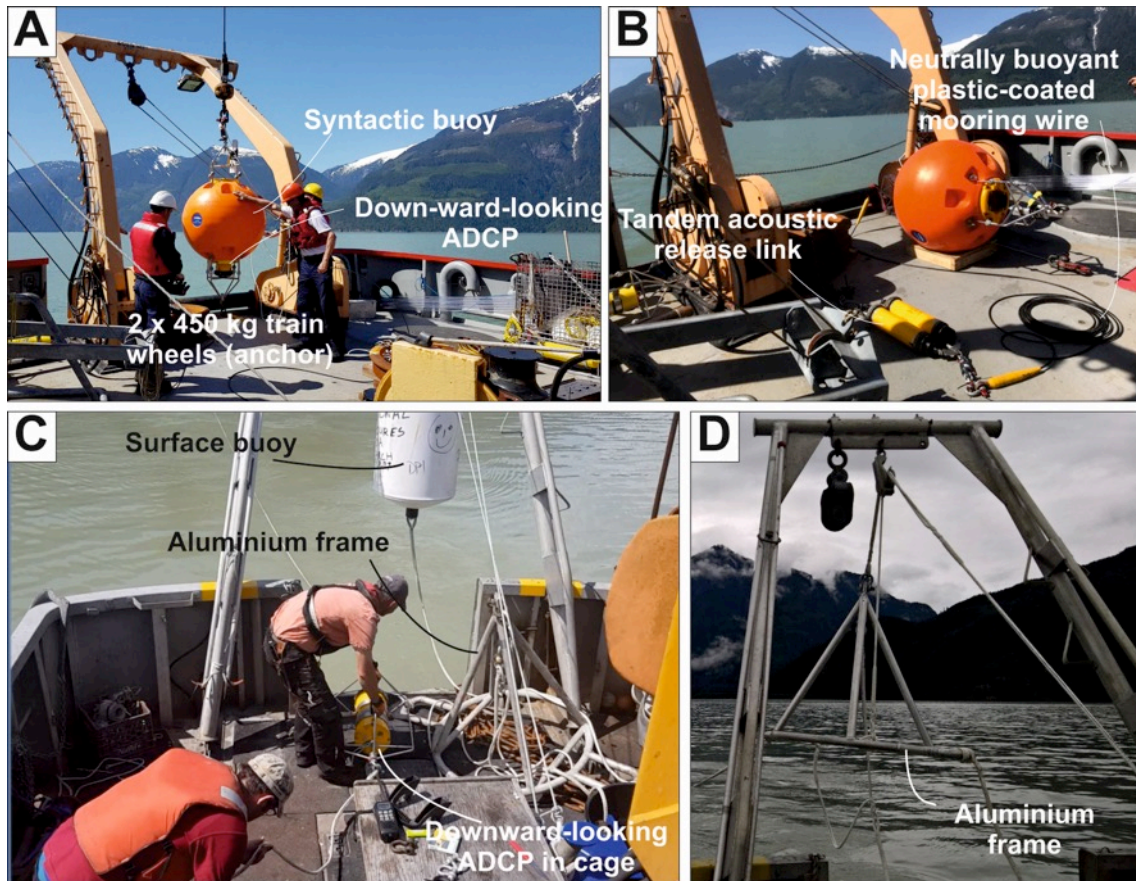
729 For narrow channels, the greater the height, the higher the likelihood that one of the four ADCP
730 beams will impinge on the channel flanks and thereby obscure details in the lower layers. In
731 Squamish, the ADCP height varied from 10 m to 15 m above the seabed at low tide as the high
732 shear part of the flows is significantly thinner than this. This is compounded by the fact that
733 turbidity currents are most likely at low tides (4 m range) at Squamish when elevation is lowest
734 (Hughes Clarke et al., 2012; Clare et al., 2016; Hage et al., 2019). In Bute Inlet, the ADCP
735 height was set at about 20 m above the channel base. The distance to the seafloor strongly
736 varied depending on the tide, and data from a local tide gauge had to be used to extract the tidal
737 signature from the data. These issues should clearly be borne in mind in tide-affected
738 (particularly macro-tidal) settings.

739 An unexpected phenomenon has been noted twice when ADCPs were suspended above a
740 turbidity current: the instrument package has been 'sucked down' into the flow. As the surface
741 buoy only has an excess of about 50 kg of flotation, it appears that the highest velocity flows
742 have enough turbulence to drag the ADCP frame down-stream, and ends up pulling the surface
743 buoy underwater because of limited anchoring; thus entraining the instrument into the flow. In
744 2017, the package was dragged down onto the seabed, where it sat for 30 minutes before rising.
745 In 2019, the instrument package was dragged down and buried, this time without release.
746 Fortunately the flotation was visible in multibeam water column imaging (110 m below the
747 surface) and could be reached with a grapple.



748

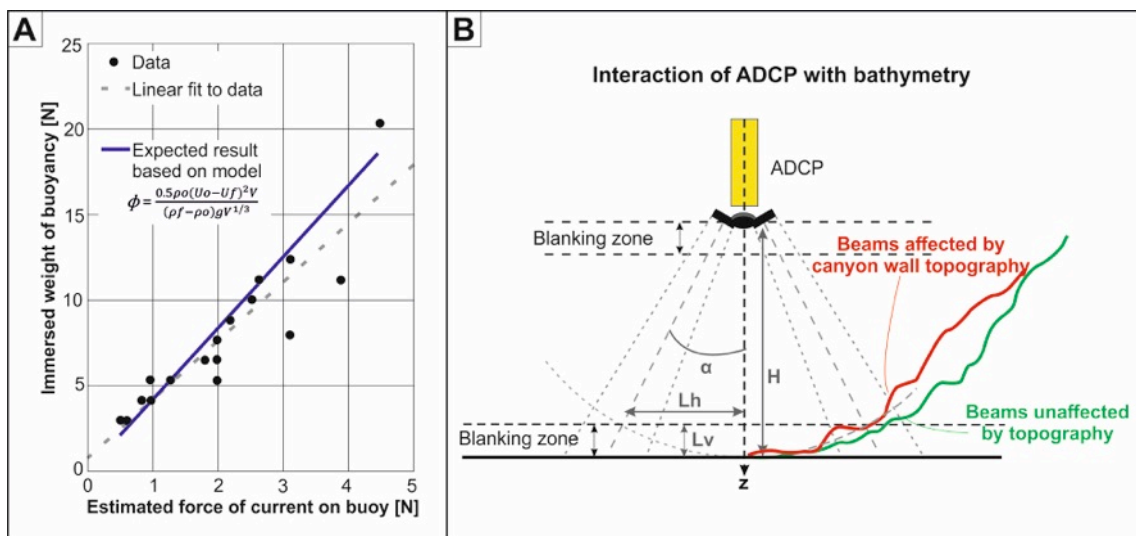
749 **Figure 10: Schematics illustrating Squamish experimental set-up from 2013 (A-C) and**
 750 **2017-2019 mooring (D). Vessel shown in blue. Acoustic imaging coverage shown in green.**
 751 **A: Location of the four anchors (all located outside the active channel areas) and all**
 752 **acoustic imaging coverage, relative to the delta lip and prodelta channels. B: Showing**
 753 **details of the offset surface buoys that allowed for azimuth stability of the suspended**
 754 **sonars. C: View from the delta lip which was 300m away, illustrating the geometry**
 755 **relative to the triggering mechanisms upstream. D: Location of the channel mouth two**
 756 **point anchor mooring in 2019. The water depths at the mooring were ~160 m on the floor**
 757 **of the channel which is about 80 m wide and 5 m deep at that point. No vertical**
 758 **exaggeration.**



759

760 **Figure 11: Photographs comparing typical hardware for single-point moorings (A & B)**
 761 **with hardware required for two-point moorings (C & D). Not shown are the 50 kg grab**
 762 **anchors used to secure each of the lines for two-point moorings.**

763



764

765 **Figure 12: A) Laboratory test of a theoretical model to determine force of current exerted**
 766 **on different weights of buoyancy. B) Schematic to show configuration of ADCP beams,**

767 **how interaction with a topographically variable seafloor may affect data quality and how**
768 **the height of the ADCP affects the proximity to seafloor at which currents can still be**
769 **monitored (see also Table 2).**

770

771 **4.3. Seafloor platforms and cabled observatories**

772 While there are clear benefits in the deployment of autonomous monitoring platforms, such as
773 moorings, they currently have finite battery power and data memory (which in turn limit
774 sampling frequency). To measure power, turbulence and fine structures within flows at high
775 temporal and vertical resolution, high bandwidth data are necessary. This may be possible for
776 moored systems when reliable methods of reconditional sampling can be developed, to record at
777 high bandwidth only during turbidity currents; however, research is still required in this area.
778 Experience at the Fraser Delta has demonstrated that it is possible to design a cabled seafloor
779 observatory that is capable of withstanding powerful turbidity currents and can transmit data in
780 real time (Lintern et al, 2019). In many settings, such as the Fraser Delta, turbidity currents
781 occur only at certain times of the year, and extreme flows may occur years apart. Capturing
782 these events at high bandwidth and long intervals apart is impossible with battery powered
783 instruments. Cabled installations provide both power, and the highest bandwidth, to a number of
784 instruments. Cabled instruments report live to shore; hence event detection is possible, which
785 might enable a response to investigate conditions shortly after the event (as was the case of
786 Lintern et al., 2016). Due to the array of instruments on the network, the exact environmental
787 conditions under which turbidity currents occur are well understood at the Fraser Delta (strong
788 freshet combined with spring tides), and their onset can be reliably predicted (Lintern et al.,
789 2019).

790 A large cabled observatory requires frequent servicing, and with current technologies can only
791 be laid with long-term dedicated resources. An advantage is that, once in operation, a scientist
792 can be assured that site visits and platform maintenance and improvement can be done
793 regularly. As mentioned, cabling platforms on the seafloor is a very expensive and intensive
794 undertaking, cannot be readily combined with other systems (unlike more mobile mooring
795 systems). There are only a few organisations worldwide currently able to maintain such a
796 system. Furthermore, the cables that provide power and distribute the data gathered are weak
797 points and are susceptible to rupture by turbidity currents (Carter et al., 2014; Clare et al., 2017).
798 Therefore serious consideration should be given to the routing of cable paths and one should
799 also be prepared for the cables to be severed. Currently, ROVs are used to connect cables to

800 platforms. This extends the deployment time from perhaps as little as a few hours on station to a
801 day or two on station, depending on tide and visibility conditions.

802 Design of seafloor platforms to monitor turbidity currents will necessarily be different from
803 more conventional tripods or other frames that are designed to measure clear-water flows (e.g.
804 Cacchione et al., 2006). Lessons learned from the Fraser Delta deployment are similar to those
805 for single-point moorings. The design challenge is to strike a balance between reducing the
806 surface area of the platform to reduce drag and increasing the weight of the structure or type of
807 legs, and to ensure it is stable to withstand toppling. For instance, the final, and most successful
808 design to date at the Fraser Delta, has the largest surface area, and has been stabilised in other
809 ways. Lessons learned for the deployment of seafloor platforms therefore include:

- 810 1) Heavy weight (e.g. 900 kg) beneath the platform, which is released when it comes
811 to retrieve the platform.
- 812 2) Stable design (e.g. tripod or quadrupedal frame) with legs that can penetrate into the
813 seafloor to act as mini piled foundations. If feet or legs are likely to become
814 embedded or buried, they should be released during retrieval. Where the feet are not
815 removable, the solution to recovering a buried platform is not to winch the platform
816 out of the sediment, but instead to apply tension and let the recovering ship slowly
817 rock the platform free.
- 818 3) Where instruments need to be suspended on hanging arms, the frame should be
819 designed such that they can be deployed at seafloor by an ROV, to reduce the
820 amount of deck space needed, and to minimise the risk of damage during
821 deployment.
- 822 4) Mounting of instruments should be reinforced and use lightweight, durable
823 materials such as titanium. Various mechanisms (hinged arms, telescoping poles)
824 may be used to extend instruments away from the platform-induced vortices,
825 towards the upstream flow to trigger other instruments. It may be appropriate to
826 consider housing instruments such as ADCPs or hydrophones in shrouded cages to
827 minimise environmental noise and vibrations. The ADCPs on the Fraser Delta
828 frame were set in a dual-axis stabilised gimbal, which righted itself and continued
829 to measure flows, even when the platform was completely upside down.

830

831 **4.4. Placement of moorings and seafloor platforms**

832 Given the efforts to ensure that monitoring platforms can successfully withstand and measure
833 turbidity currents, it would be unfortunate if they were not deployed in the correct location.

834 Precise placement also remains a challenge, particularly where support from ROVs (i.e. to
835 verify placement location or assist with re-siting) is unavailable, or is considered too time-
836 consuming or costly. A high quality base map is essential to ensure the proposed target is
837 appropriate. As the seafloor elevation and planform can vary considerably in active submarine
838 canyons and channels (e.g. Paull et al., 2018; Gales et al., 2019; Vendettuoli et al., 2019;), it is
839 recommended that multibeam bathymetric data be acquired prior to deployment to accurately
840 determine the water depth and seafloor relief to ensure that the proposed location is correct (e.g.
841 the canyon thalweg has not migrated, ADCPs will not be affected by interference with canyon
842 side walls, mooring is not placed on a canyon-wall slump etc).

843

844 ***4.4.1. Deployment and siting of moorings***

845 When placing moorings in submarine canyons or channels, the desired seafloor targets are
846 usually very small and may rely on deployment from vessels without dynamic positioning (a
847 computer-controlled system to maintain position and heading using thrusters). Thus, the vessel
848 may drift off location easily during the deployment. Even with dynamic positioning, moorings
849 dropped from the sea surface can drift with the current or during free fall. A triangulated
850 location is typically acquired for moorings by communicating with the acoustic releases;
851 however, this is often inaccurate, difficult in great water depths, and can be complicated by
852 echos from steep-sided canyon walls or other topographic features.

853 Another option to determine the location of moorings is to make use of a multibeam
854 echosounder. As long as the mooring array has a series of scattering targets (flotation spheres or
855 instrument housings) that are separated by more than the sonar range resolution, they can
856 usually be discerned from the natural scatterers, as you pass over them. This method has been
857 used for detecting location of moorings, as well as to image passing turbidity currents (Hughes
858 Clarke et al., 2014), in shallow water fjord settings, and is also feasible in deeper water using
859 the multibeam system constrained to shorter pulse lengths (2 ms) in a narrow swath. This should
860 therefore enable identification of moorings in up to 2 km of water.

861 Where moorings are lowered to seafloor, a position fix can be acquired from an ultra-short
862 baseline (USBL) system. It is worth including beacons on the moorings that would allow the
863 mooring's actual position (during deployment and monitoring periods) to be determined with
864 the necessary accuracy; however, this technique gets increasingly expensive with greater water
865 depths.

866 The effects of human interference with the seafloor should be considered when choosing a
867 platform location, as , activities such as fishing, trawling, anchor deployment and dredging can
868 snag, displace or damage monitoring platforms. Moorings should be placed in water depths
869 greater than the keel of icebergs in areas affected by seasonal ice cover.

870

871 ***4.4.1. Specifics on deployment of single-point moorings***

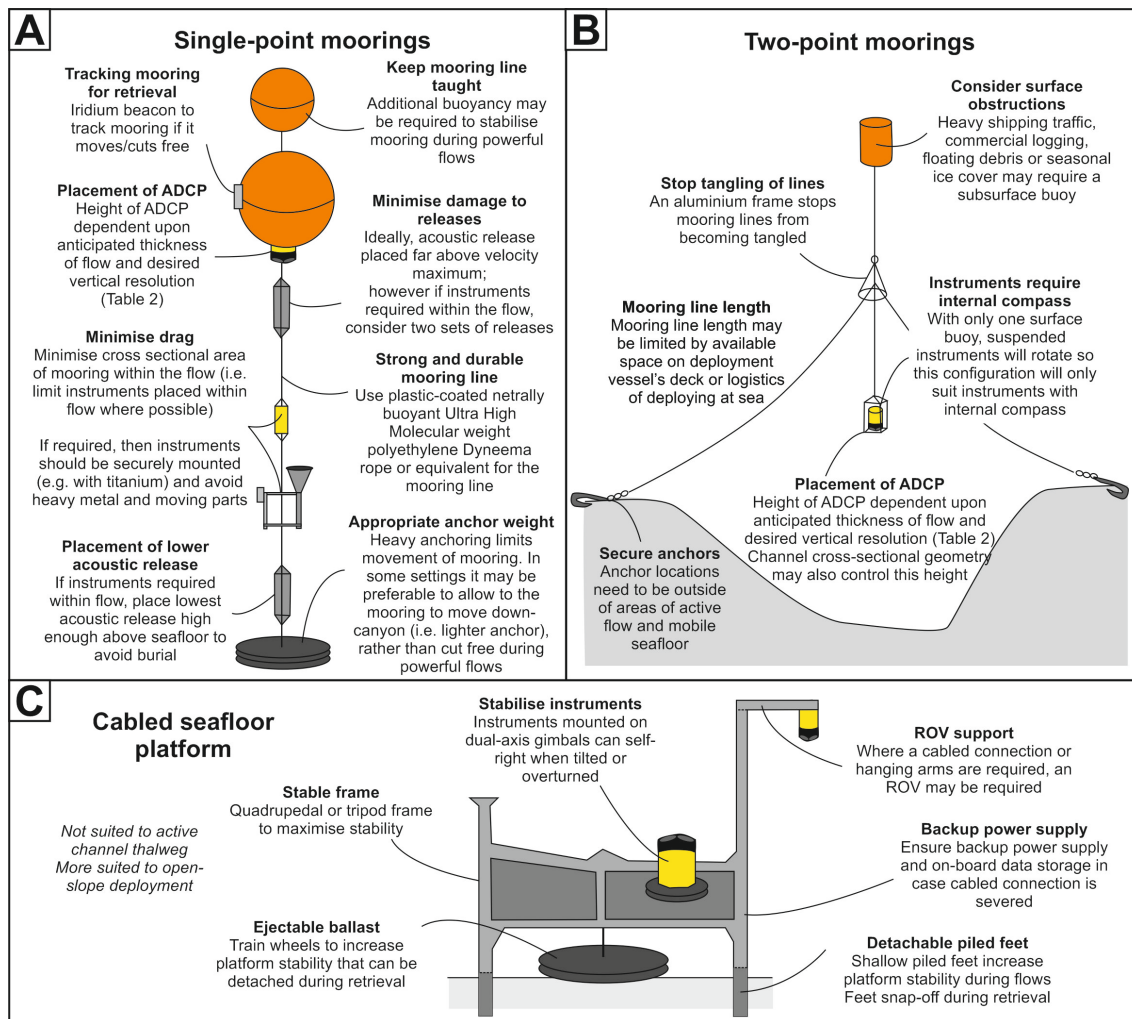
872 Two general approaches exist for the deployment of single-point moorings. The first is to
873 deploy the anchor last (i.e. buoyancy and mooring line with instruments attached are offloaded
874 to sea prior to dropping the anchor at the desired location). An anchor-last deployment also
875 allows you to manoeuvre the vessel to above the desired location using USBL, and then drop
876 the mooring once on location. This approach has been shown to achieve a precision of +/-10-20
877 m horizontal accuracy in water depths of up to 2 km, and 50-60 m at 5 km water depth , and
878 depends firmly upon the vessel's captain, ship handling skills of the mate on watch, maintaining
879 efficient communication between the Deck, Bridge and Science crew, and fair weather
880 conditions and sea state at the time of deployment. The second is to deploy the anchor first,
881 which can be hazardous as the mooring line will be in tension on the back deck of the vessel.
882 For this reason in particular, an anchor-first strategy is precluded when heavy anchors are
883 required (due to very high line tensions).

884

885 ***4.4.2. Deployment of two-point moorings***

886 In shallow water, where the line suspending the instrument and anchor lines are all connected to
887 a surface buoy, anchors for two-point moorings can first be placed individually. After the
888 anchors are placed with a small surface float, then the anchor lines can be connected to a single
889 point above the channel, and the instrument can be lowered from this central surface buoy. In
890 deeper water, the use of submerged frames is more appropriate (given the length of mooring
891 lines required). Two deployment methods have been successful in safely placing these deeper
892 water two-point moorings. In the first method, the instrument was lowered above the channel,
893 followed by the frame, and roughly kept in place by a small boat. While the small boat held on
894 to the second anchor line, a larger ship (with winch and A-frame) sets off with the first anchor
895 line. On the larger boat, the anchor line is connected to the chain and anchor before being
896 dropped at the anchor location. Then the larger ship returns to the smaller boat to pick up the
897 second anchor line and drop the second anchor.

898 A second approach, that has also been successfully applied, involves releasing the central part of
899 the mooring down in one step. For this approach, all the lines, the instrument, the frame and the
900 float need to be carefully laid out on the back deck. The procedure starts by deploying the first
901 anchor and laying out the anchor line, while the ship slowly steams from the first anchor
902 position towards the channel. Just before reaching the channel, the instrument is lowered into
903 the water with a line tied to the submerged frame that is hanging from the A-frame of the ship.
904 As the boat crosses the channel, the first anchor line starts to tighten and the submerged frame is
905 dropped in the water. While the boat keeps steaming slowly towards the second anchor position,
906 the second anchor line and the line connecting the submerged frame to the surface buoy are
907 slowly released. As the line towards the buoy runs out, the buoy is released from the ship.
908 Finally, when the boat reaches the second anchor position the last anchor is dropped. The
909 advantage of the first method is that the deployment is done step-by-step and is more controlled;
910 however, it requires two vessels and there is a higher chance for the instrument and second
911 anchor line to become tangled during the deployment of the first anchor. The second method
912 requires only one vessel, but needs a larger back deck (and very careful preparations), as the
913 ship dragged the lines behind the ship and its propeller, the ship will have to continue moving
914 forwards to prevent the lines from tangling in the propeller. So in the second approach
915 everything needs to be deployed in one go; once the first anchor is dropped there is no way
916 back. Both methods have been in Bute Inlet four times, and all moorings have been placed
917 successfully. Retrieving these two-point moorings is fairly straightforward. After picking up the
918 surface buoy the line is connected to the winch and the whole mooring is pulled out. A 1 tonne
919 winch has always been successful in retrieving the moorings, although we have had to cut one
920 anchor line, possibly as a result of a buried anchor. Depending on the type of anchor and the
921 angle of the anchor lines, larger forces could be applied to the submerged frame, so it might be
922 advisable to make sure that the link between the anchor lines and the frame are the weakest
923 connection in the mooring, to ensure that the instrument is always recovered. Alternatively,
924 acoustic release links could be incorporated into the mooring design (i.e. one on each mooring
925 line); however, these would add additional weight to the mooring line which would need to be
926 considered.



927

928

Figure 13: Summary of lessons learned for designing monitoring platforms, illustrating key considerations when measuring powerful turbidity currents.

929

930

931

5. Conclusions and final thoughts

932

The design of monitoring platforms needs to deal with high velocities and sediment

933

concentrations close to the seafloor, capable of tilting, displacing, transporting and even

934

damaging instruments. Our experience shows that, despite the challenges posed it is possible to

935

make detailed measurements of powerful sediment-laden flows. These challenges can be

936

overcome by simplifying single-point mooring design to reduce drag potential, or deploying

937

two-point moorings (or from vessels), where neither mooring lines nor instruments interact with

938

the flow itself. Where it is necessary to deploy instruments within the flow, it may not be

939

possible to reduce drag, hence additional stability is essential, such as extra buoyancy and

940

anchoring for single-point moorings, or piled legs and extra weight for seafloor platforms.

941

Instrument mounting may be a weak point in such designs; hence brackets and cages should be

942 more robust than for standard moorings. Table 3 provides a summary of these considerations
943 and scenarios that are most suitable for different monitoring platforms.

944

945 There is currently a push to develop next generation monitoring tools to detect and characterise
946 turbidity currents; relying upon passive detection, rather than direct measurements (e.g. Clare et
947 al., 2017; Lintern et al., 2019). Such tools include hydrophones and geophones and will enable
948 measurement of turbidity currents, and other submarine mass movements, without the need to
949 place moorings or platforms in the path of the flow (particularly where the flow is restricted to
950 channels (Chadwick et al., 2012; Caplan-Auerbach et al., 2014). This approach requires
951 calibration against ADCPs and other measurements, and initial results are promising. There is
952 clear evidence that acoustic signals can be linked to independently-measured turbidity currents
953 (Hatcher, 2017; Lintern et al., 2019). In addition to measuring transit speeds via arrival times,
954 there is potential to measure some basic features of flow character using hydrophones. For
955 example, the intensity of acoustic signals may be related to internal flow speeds (via intensity of
956 grain collisions), grain size (sand or mud dominated flow) or the presence of a dense and coarse
957 near-bed layer. However, further work would be needed to determine what is possible, and how
958 flows are recorded. Other developments in distributed sensing along fibre-optic cables also
959 demonstrate the potential utility of cabled submarine links, such as those that connect the Fraser
960 Delta Dynamics Laboratory to the VENUS seafloor cabled network, to measure strain,
961 temperature and to use the optical fibres as distributed acoustic sensors (e.g. Lindsey et al.,
962 2017, 2019; Hartog et al., 2018).

963

964 Finally, there is growing interest in monitoring a wider range range of deep-sea sediment
965 transport processes, including the influence of internal tides (Maier et al., 2019b), thermohaline-
966 driven circulation (Miramontes et al., 2019), and the mixed interaction of down-slope gravity-
967 driven flows such as turbidity currents with along-slope contour currents (Normandeau et al.,
968 2019b). As such flows are typically of lower velocity (generally $\ll 1$ m/s; McCave et al., 2017)
969 and comprise lower sediment concentrations than turbidity currents, they should be considerably
970 more straightforward to measure. Therefore many of the issues outlined in this study are
971 unlikely to be a major issue; however, the lessons learned should still be considered – such as
972 minimising drag and maintaining stability of the platform to ensure that high quality results are
973 acquired. Burial risk may be greater in areas of high net deposition. To date, limited near-bed
974 measurements of contour currents have been made, and none are yet known from mixed
975 turbidity current-contour current systems. Therefore, there is a need for instruments to be placed
976 closer to seafloor in such systems to fill this knowledge gap. Such systems are also typically

977 much more laterally extensive than “conventional” turbidity current canyons or channels; hence
978 it will be necessary to deploy an array of monitoring platforms to characterise the spatial
979 variability in near-bed flow that may be strongly controlled by local variations in seafloor
980 morphology.

981

982 We conclude that recent and ongoing advances in technology and mooring design will ensure
983 that key knowledge gaps in turbidity current behaviour can soon be filled, providing valuable
984 information for designing resilient seafloor infrastructure, and understanding of how and when
985 these globally important processes transport sediment, nutrients and organic carbon to the deep
986 sea.

987

988 **Acknowledgements**

989 This study is based on fieldwork and logistical support provided by a large number of people,
990 including: i) the Monterey Coordinated Canyon Experiment (CCE) Team; ii) Bute Inlet 2016
991 and 2018 teams; iii) Squamish 2015 monitoring team; and most notably, iv) the patient and
992 diligent crews of the following vessels: Strickland, Heron, Vector, Rachel Carson and Dock
993 Ricketts. Clare, Cartigny, and Talling acknowledge funding from the Natural Environment
994 Research Council (NERC), including “Environmental Risks to Infrastructure: Identifying and
995 Filling the Gaps” (NE/P005780/1) and “New field-scale calibration of turbidity current impact
996 modelling” (NE/P009190/1). The authors also acknowledge discussions with collaborators as
997 part of Talling’s NERC International Opportunities Fund grant (NE/M017540/1) “Coordinating
998 and pump-priming international efforts for direct monitoring of active turbidity currents at
999 global test sites”. Talling was supported by a NERC and Royal Society Industry Fellowship
1000 hosted by the International Cable Protection Committee. Funding for the Coordinated Canyon
1001 Experiment (CCE) was provided by David and Lucile Packard Foundation, Natural
1002 Environment Research Council (grant NE/K011480/1), U.S. Geological Survey (USGS) Coastal
1003 and Marine Program, and Ocean University of China. Additional funding for MAC was
1004 provided by NERC National Capability Climate Linked Atlantic Sector Science Programme
1005 (NE/R015953/1).

1006

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1319 **Table 1: Examples of adverse effects to monitoring platforms and instruments caused by**
 1320 **turbidity currents from previous studies and sites referenced in this paper.**

Location and literature source	Instrument type	Water depth	Maximum measured velocity	Documented adverse effects, damage etc.
Scripps Canyon (Inman, 1970, Inman et al., 1976)	Point current meter 4 m above seafloor connected to shore by a cable	46 m	1.9 m/s	Current meter failed during flow and was subsequently lost. Flows bent a 2.5 cm-thick solid steep rod bolted into canyon bedrock
Lake Geneva, Switzerland (Lambert & Giovanali, 1988)	Point current meters on single point moorings	Up to 170 m	Up to 3m/s	Anchor cables broke and moorings floated to surface
La Jolla Canyon, California (Shepard and Marshall, 1973)	Point current meters on single point mooring	200 m	Up to 0.5 m/s before data recording stopped	Moorings displaced 500 m down-canyon
Open slope, Hawaii (Dengler et al., 1984)	Point current meters on single-point mooring	Up to 600 m	Up to 2 m/s	Episodic down-slope movement of moorings by 2.4 km
Squamish Delta, British Columbia (Hughes Clarke et al., 2009, 2012)	Upward-facing ADCP mounted in seabed frame	Up to 150 m	Up to 1.5 m/s	ADCP frame buried by 2 m of sediment.
Fraser Delta, British Columbia (Lintern et al., 2016)	Cabled seafloor frame (1 tonne) fitted with numerous instruments including upward-facing ADCP	40-107 m	Up to 10 m/s	Platform tumbled down delta and severed connection with onshore cable
Bute Inlet, British Columbia (Prior et al., 1987)	Point current meters, Anderson-style sediment traps on single point moorings. Seafloor frame vane deflectors.	Up to 520 m	Up to 3.4 m/s	Rotors and vanes on current meters broken off or fouled (causing poor data quality), shackles and stainless steel frames bent and sheared, some entire instruments lost. Mooring wires parted, releasing instruments to surface. Moorings displaced along- and down-channel (up to 1 km). Acoustic releases failed to detach due to assumed burial by sand.
Monterey Canyon, California (Paull et al., 2002)	Seafloor trapezoidal frame (97 cm by 83-cm base and 48 cm tall)	525 m	N/A	Frame transported 550 m down-canyon and buried in up to 0.7 m of sediment.
Monterey Canyon California (Paull et al., 2010)	Trawl resistant seafloor frames (up to 1360 kg)	289 and 520 m	N/A	Moved up to 170 m down-canyon and buried in up to 1.5 m sediment.

Monterey Canyon, California (Paull et al., 2018)	Array of single-point moorings, a seafloor frame, and a 800 kg frame carrying a transponder	Up to 1,850 m	Up to 7.2 m/s	MS1 transported 7.1 km down canyon before breaking loose and floating to surface, sediment traps torn apart, 800 kg frame transport 4.5 km down canyon and buried in >1 m sand.
Congo Canyon, West Africa (Khripounoff et al., 2004; Vangreisheim et al., 2009)	Point current meters on single-point mooring	Up to 4,790	Up to 3.5 m/s	Tilting of mooring prior to parting of mooring anchor line, releasing instruments to surface. Damaged current meter (30 m above seafloor) and sediment trap (40 m above seafloor)
Congo Canyon, West Africa (Cooper et al., 2012; Azpiroz-Zabala et al., 2017a)	Down-ward facing ADCP on single-point mooring	2,000 m	Up to 2.5 m/s	Rotating ADCP Interference with canyon sidewall

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1322 **Table 2: Seafloor footprints of ADCP beams for different vertical heights, assuming a 20**
 1323 **degree beam angle (typical for the moorings discussed in this paper). Illustrated in Figure**
 1324 **12.**

Height above seafloor, H [m]	Side lobe interference zone – blanked above seafloor, L _v [m]	Radius of ADCP beam footprint at seafloor, L _h [m]	Diameter of ADCP beam footprint at seafloor, 2 x L _h [m]	Example ADCP frequency as discussed in this paper
300	18.1	102.6	205.2	Var Canyon 75 kHz ADCP
85	5.1	29.1	58.1	Congo Canyon 300 kHz ADCP
70	4.2	23.9	47.9	Monterey Canyon 300 kHz ADCP
35	2.1	12.0	23.9	Bute Inlet 600 kHz ADCP
12	0.7	4.1	8.2	Squamish 1200 kHz ADCP

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1326 **Table 3: Summary of considerations for different types of turbidity current monitoring**
 1327 **platforms**

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Platform type	Environment	Benefits	Considerations
Single-point mooring	Long-term deployment in canyon/channel thalweg or unconfined slope	<ul style="list-style-type: none"> • Simple mooring design • Relatively easy to deploy • Simple retrieval using acoustic release link and sacrificial anchor weight 	<ul style="list-style-type: none"> • As it interacts with the flow, mooring may need to be designed to cope with down-slope transport or maintain taught line (large anchor weight and high buoyancy) • Ideally, reduce drag by minimising cross sectional area (e.g. reducing instruments) on mooring line

			<ul style="list-style-type: none"> • Acoustic releases should be placed above velocity maximum of flow; however if instruments are required within the flow, then releases should be placed below those instruments and also above • Large ocean-going vessel may be required to deploy heavy anchor weights and buoyancy
Two-point mooring	<p>Short-term deployment over channels/canyons</p> <p>Particularly useful where flows are highly erosive or have dense near-bed layer</p>	<ul style="list-style-type: none"> • None of mooring interacts with flow • Unaffected if erosion or deposition affect seafloor 	<ul style="list-style-type: none"> • Challenging field deployment requiring considerable lengths of mooring line • Requires larger vessel for retrieval of anchors and mooring lines • Only possible where stable terraces, levees or channel margin permit anchor placement • Surface buoy may pose a problem in areas with seasonal ice cover, busy shipping or logging
Vessel-mounted mooring	Shallow water settings where timing of turbidity currents is known	<ul style="list-style-type: none"> • None of mooring interacts with flow • Continuous power to instruments • Possible to adjust instrument settings and acquire calibration samples in real-time 	<ul style="list-style-type: none"> • Only suitable for shallow water settings • Requires crewed vessel; hence, only suitable for relatively short deployments
Benthic lander	Unconfined slope and/or dilute flows	<ul style="list-style-type: none"> • Continuous data transmission, enabling near-real time response (e.g. to perform seafloor survey) • Externally powered, allowing for multiple instruments recording at high frequency 	<ul style="list-style-type: none"> • Not suitable for placement in active canyon/channel thalweg • May be buried, or undermined by erosion • To withstand powerful flows, requires ejectable ballast and removable feet • Requires support ROV during deployment if cabled links or additional instruments need to be added once platform is on seafloor

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