

1 **Diffusion of Halogens (F, Cl, Br, I) in silicic melt**

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36

37 **Abstract**

38 *Chemical diffusion of the halogens F, Cl, Br, and I in silica-rich natural melts was*
39 *experimentally investigated by the diffusion couple technique. Experiments were conducted*
40 *under anhydrous conditions at atmospheric pressure and hydrous conditions (~1.5 wt.%*
41 *H₂O) at 160 MPa, over a temperature range of 750–1000 °C and 1000–1200 °C, respectively.*
42 *Quenched trachytic melt samples were analyzed using an electron microprobe (EPMA) and*
43 *secondary ion mass spectrometry (SIMS).*

44 *All halogens exhibit Arrhenian behavior during diffusion in the investigated melt*
45 *compositions with F always diffusing fastest. The other halogens show progressively slower*
46 *diffusion (F > Cl > Br > I) correlated to their ionic radii. In anhydrous melt a diffusivity range*
47 *of 3–4 orders of magnitude is covered among the halogens with $D_F(1000\text{ °C}) \sim 5 \times 10^{-13}\text{ m}^2\text{s}^{-1}$*
48 *and $D_I(1000\text{ °C}) \sim 1 \times 10^{-16}\text{ m}^2\text{s}^{-1}$. The diffusivities of all halogens increase in hydrous melt*
49 *yielding for example $D_F(1000\text{ °C}) \sim 3 \times 10^{-12}\text{ m}^2\text{s}^{-1}$. However, the largest increase is observed*
50 *for the slowest-diffusing halogens, resulting in a narrower overall diffusive range of only 1–*
51 *2 orders of magnitude with iodine diffusivity yielding $D_I(1000\text{ °C}) \sim 9 \times 10^{-15}\text{ m}^2\text{s}^{-1}$. Activation*
52 *energies (E_A) of all halogens consistently range from ~200–290 kJ mol⁻¹ in anhydrous melts.*

53 *In hydrous melt E_A generally decreases, with the highest decrease determined for F (~131 kJ*
54 *mol⁻¹) and only slight changes for the other halogens (~201–222 kJ mol⁻¹).*
55 *Our diffusivity data of the anhydrous series exhibit a pronounced correlation of diffusivity*
56 *with the ionic radii, suggesting that halogen diffusion in highly polymerized melt is closely*
57 *related to the melt's ionic porosity. The correlation between diffusivity and ionic radius is*
58 *only weakly observed in the hydrous experiments indicating that the ionic porosity is*
59 *sufficiently large to weaken the rate-limiting effect of the ionic radius due to the more*
60 *depolymerized melt structure in the hydrous case. In hydrous experiments, the process of*
61 *ionic detachment becomes more important as a rate-limiting diffusion mechanism,*
62 *comparable to the case of diffusion of divalent/trivalent cations or halogen diffusion in*
63 *basaltic melt.*
64 *The results of this study provide the first consistent diffusion dataset including all halogens*
65 *under naturally relevant magmatic conditions and highlight the pronounced compositional*
66 *effect of both, major element and dissolved H₂O on halogen diffusion. These data emphasize*
67 *the potential of diffusive fractionation among the halogens, especially in a melt of low water*
68 *content, which may be applied as a monitoring tool for volcanic unrest on actively degassing*
69 *volcanoes.*

70 **Keywords:** *halogens (F, Cl, Br, I); diffusion couple; ionic porosity; iodine; silicate melt*

71

72 **1. Introduction**

73 Volatiles are among the most influential constituents of natural silicate magmas. They
74 affect various chemical and rheological melt parameters, which in turn have crucial
75 implications for individual volcanic eruption behavior. Typically, volcanic volatiles are
76 dominated by water (H₂O) and carbon dioxide (CO₂), however, sulfur-species (e.g., SO₂)
77 and halogens, mainly F and Cl, can reach significant concentrations (up to few wt.%) in

78 silicic melts and volcanic gases (e.g., Aiuppa et al., 2009; Dolejš and Zajacz, 2018).
79 Halogens are known to exert strong effects on the physico-chemical properties of silicate
80 melts such as melt viscosity or phase equilibria (e.g., Manning, 1981; Mysen and Virgo,
81 1985; Webster et al., 2018). For example, F is known to reduce the viscosity of silicic melts
82 to a similar degree as added H₂O (Dingwell et al., 1985; Dingwell and Hess, 1998; Zimova
83 and Webb, 2007; Giordano et al., 2008; Baasner et al., 2013), and can depress the isobaric
84 liquidus temperature in basalt (e.g., Filiberto et al., 2012) and felsic melts (e.g., Dolejš and
85 Baker, 2007b; Dolejš and Baker, 2007a). Even though more complex, substantial effects
86 on the liquidus temperature have also been reported for Cl-bearing natural melts of both
87 basaltic (e.g., Filiberto et al., 2014; Farcy et al., 2016) and silicic systems (Feisel et al.,
88 2022). The physico-chemical effects of added Br and I to a melt are, by contrast, far less
89 understood. Concentrations of Br and I are typically lower than those of F and Cl,
90 consistent with an increasing fluid-melt partition coefficient with increasing ionic radius
91 of Cl, Br and I (181, 196 and 220 pm, respectively; Bureau et al., 2000). F (133 pm)
92 instead has an ionic radius very similar to O²⁻ and OH⁻ (140 and 137 pm; Shannon, 1976)
93 and can therefore easily substitute for these components which promotes its solubility
94 and compatibility in the melt (Balcone-Boissard et al., 2010).

95 Volcanic halogen degassing can have significant impacts on the environment and Earth's
96 climate, for example by the depletion of stratospheric ozone by Cl and Br (Bobrowski et
97 al., 2003; Bobrowski et al., 2007; von Glasow et al., 2009; Surl et al., 2015; Roberts, 2018).

98 The degassing of F and Cl into isolated pores of slowly cooling magma was shown to cause
99 etching of silica and re-precipitation of vapor-phase cristobalite (e.g., Horwell et al., 2010;
100 Schipper et al., 2020).

101 Due to the high impacts that halogens can have on volcanological and environmental
102 processes, they have received increasing attention in recent years and are routinely

103 measured in volcanic plumes. Halogen exsolution was shown to occur at relatively
104 shallow depths compared to the release of H₂O or CO₂ (e.g., Spilliaert et al., 2006; Schipper
105 et al., 2019) and may therefore help to characterize volcanic degassing. Due to the large
106 differences in diffusivity between water and the halogens, as well as amongst halogens
107 themselves, diffusive fractionation is expected to occur during melt inclusion entrapment
108 or bubble and crystal growth (e.g., Alletti et al., 2007). Relative halogen (or halogen
109 ratios) measurements in volcanic gases and plumes could therefore be applied to track
110 diffusive fractionation processes in order to monitor volcanic activity at depth.

111 Even though halogen diffusion has been investigated since the 1980s our knowledge is
112 still far from complete. It was shown that halogen diffusion is independent of its
113 concentration and follows Arrhenian behavior in silicate melts. A weak pressure-
114 dependence on diffusivities was reported by some authors (e.g., Dingwell and Scarfe,
115 1984; Bai and Koster van Groos, 1994), however, most recent studies have found only
116 minor pressure effects (e.g., Alletti et al., 2007; Balcone-Boissard et al., 2020). Many
117 studies focused on simplified model systems (Dingwell and Scarfe, 1984; Dingwell and
118 Scarfe, 1985; Bai and Koster van Groos, 1994) and most investigated F and/or Cl only
119 (Balcone-Boissard et al., 2009; Böhm and Schmidt, 2013; Fortin et al., 2017; Yoshimura,
120 2018; Feisel et al., 2019). To date only two studies include Br diffusion (Alletti et al., 2007;
121 Balcone-Boissard et al., 2020) and, to our knowledge, no study has addressed iodine
122 diffusion.

123 In this study we applied the diffusion couple technique on three different homogeneous
124 melts synthesized from natural silicic pumice and lava to determine the diffusivity of all
125 halogens (F, Cl, Br, I) over a temperature range relevant to the studied magmatic systems.
126 Concentration profiles of F and Cl were measured by electron microprobe (EPMA) and
127 additional profiles were acquired by secondary ion mass spectrometry (SIMS) covering

128 all halogens. We report data for both anhydrous and hydrous (~1.5 wt.% H₂O) conditions
129 and this is — to our knowledge — the first study to report iodine diffusion data in natural
130 silicate melts. These results highly improve our understanding of magmatic halogen
131 transport mechanisms at crustal conditions and expand the database of halogen diffusion
132 to advance a thorough understanding of volcanic processes.

133

134 **2. Methods**

135 **2.1 Experimental procedure**

136 Two different starting materials were used in this study comprising glass synthesized
137 from tephra of the Hekla (Iceland) H3 eruption (e.g., Thordarson and Larsen, 2007;
138 Weber and Castro, 2017), and glass synthesized from natural obsidian lava of the 2011
139 Cordón Caulle eruption, Chile (e.g., Castro et al. 2013, 2016; Schipper et al. 2013, 2019;
140 Alloway et al. 2015). Synthesis and experimental methods utilizing the Hekla material
141 (“HX” series – “Hekla experiments”) are described in detail in the study of Feisel et al.
142 (2019). Synthesis of starting glass derived from Cordón Caulle obsidian lava (“CCX” series
143 – “Cordón Caulle experiments”) is described in detail in Feisel et al. (2022). The general
144 approach used for both melts is briefly described below.

145 The initial sample synthesis was carried out in the laboratories of the Earth Science
146 department of the LMU Munich. Natural obsidian was crushed, synthesized and
147 homogenized at high temperature (1400–1550 °C) for ~2 days to produce almost
148 completely volatile-free glass. After the first synthesis, the batch was crushed and split.
149 One half was enriched with halogen-bearing Na-salts (NaF, NaCl, NaBr, NaI) so that each
150 halogen would account for approximately 1 wt.% of the whole batch after mixing. The
151 halogen-depleted split was in turn enriched with a certain amount of Na₂CO₃ to account
152 for the sodium enrichment in the halogen-bearing glass, due to the use of aforementioned

153 Na-salts to add halogens. Each aliquot was again synthesized and homogenized at high T
154 before preparation of ca. 2 mm thick polished glass discs (4.6 mm diameter) and crushed
155 glass powder to be used in anhydrous and hydrous diffusion experiments, respectively.

156

157 *2.1.1 Anhydrous diffusion experiments*

158 Diffusion couples were constructed using 5 mm platinum tube with a wall thickness of
159 0.2 mm. The bottom of the capsule was closed with a Pt lid which was pressed onto and
160 welded to the tube using an arc-welder. Afterwards the bottom of the capsule was tamped
161 into a cylindrical shape. One halogen-enriched glass disc was loaded into the bottom of
162 the capsule and a disc of the equivalent halogen-depleted glass was loaded on top of it,
163 resulting in both pieces touching at their polished surfaces. Each capsule was closed by
164 welding a second lid to the top of the capsule, which ensured closed-system conditions.
165 The diffusion experiments were carried out in a vertical tube furnace at atmospheric
166 pressure. Each capsule was first weighed and then loaded into an alumina tube that is
167 closed on the bottom end, making sure the halogen-depleted half was facing upwards.
168 After the furnace equilibrated at the target temperature the alumina tube was inserted
169 vertically into the furnace through a small opening in the top, ensuring that the capsule
170 was sitting upright on the closed bottom of the alumina tube in the hot zone of the
171 furnace. The temperature was monitored by the internal thermocouple and an external
172 K-type thermocouple being in direct contact with the experimental capsule. Additionally,
173 most experiments were monitored using a second external K-type thermocouple
174 interfaced with an Arduino microprocessor to read the temperature remotely and store
175 the *P-T-t* data to a computer. Temperatures are estimated to be accurate to ± 2 °C based
176 on the manufacturer's (*Omega*) reported accuracy of the thermocouple device. The
177 experimental durations ranged from ~20 hours to 35 days depending on the temperature

178 investigated. The time needed for the capsule to heat to the target temperature ranged
179 between 3 and 5 minutes. The experiments were quenched by sliding the alumina tube
180 out of the hot furnace and letting it cool in air at room temperature in an upright position.
181 Typical cooling times to a temperature of 200 °C as indicated by the external
182 thermocouple were in the range of 3 to 4 minutes. After each experiment the samples
183 were weighed again to confirm closed-system conditions were maintained during the
184 experiment.

185

186 *2.1.2 H₂O-bearing diffusion experiments*

187 Hydrous experiments were carried out using CCX starting material only and are denoted
188 by “CCX-H”. The capsule preparation followed the procedure described in Feisel et al.
189 (2022) using powdered starting glass and Pt-tube with a diameter of 4 mm and a wall
190 thickness of 0.2 mm as capsule material. After welding the capsule shut on the bottom
191 and tamping into a cylindrical shape using a tight-fitting rod, ground glass powder of the
192 halogen-enriched samples was introduced until the capsule was filled about halfway.
193 During filling the powder was slightly tamped several times to avoid entrapment of air
194 and ensure a flat diffusion interface. The exact amount of added sample powder was
195 weighed using an analytical balance of 5-digit precision and the equivalent amount of 1.5
196 wt.% H₂O was added to the capsule using a microsyringe. The procedure was repeated
197 using the halogen-depleted powder and again the equivalent amount of 1.5 wt.% H₂O for
198 the top part of the capsule. Finally, the capsule was closed by welding a tight-fitting lid to
199 the open end and tamped into cylindrical shape with a pressure of ~750 kg using a
200 hydraulic press. Before an experiment, the capsules were repeatedly heated to ~120 °C
201 and weighed to confirm no weight loss. Capsules with significant weight loss were
202 discarded.

203

204 Experiments were performed in a *Deltech* vertical tube furnace fitted with a gas-
205 pressurized TZM (Tungsten Zirconium Molybdenum) cold-seal pressure vessel assembly.
206 The sample capsule is loaded into the cold-seal assembly which is connected to a
207 stainless-steel high-pressure tubing using Argon gas as a pressure medium. To mitigate
208 water-loss from inside the capsule to the Ar-atmosphere ca. 0.3 MPa of CH₄ was inserted
209 to the pressure system before pressurizing with Ar (e.g., Sisson and Grove, 1993; Szramek
210 et al., 2006; Shea and Hammer, 2013). The cold-seal assembly consists of an inner
211 autoclave made of TZM (~30 mm outer, ~6.5 mm inner diameter) and an outer sheath
212 made of Inconel (~50 mm diameter, 5 mm wall thickness), both attached to a water-
213 cooled base on the bottom. The outer sheath slides over the inner autoclave leaving ~5
214 mm between the two. This space is constantly purged with Ar during the experiment to
215 provide a non-reactive atmosphere for the inner pressurized autoclave. The inner
216 autoclave is tightened against a steel cylinder via a doubly coned and sealed small
217 Beryllium-bronze cylinder (“football”-shape) between the two. A second, lower vessel
218 made of stainless steel is attached to the bottom of the water-cooled coupling base and is
219 connected to the pressure line on the bottom with a high-pressure fitting. The sample
220 capsule sits in a small cup which in turn is attached to an Inconel rod which is inserted
221 into the bottom of the steel cylinder. A magnet at the bottom of the rod allows for the
222 control of the position of the sample capsule in the assembly from the outside using a
223 neodymium ring-magnet that slides over the lower steel cylinder. The whole assembly
224 including the experimental capsule is inserted into the vertical tube furnace from the
225 bottom before it is brought up to the target temperature and pressure. The pressure is
226 generated by an HIP GB-60 gas booster device operated with compressed air and
227 conveying Ar gas as a pressurizing medium.

228 During the time of heating, the capsule remains in the water-cooled part of the assembly
229 (rapid-quench coupler) which is outside the furnace and constantly flushed with water
230 (~15 l/min). When the target P - T conditions are reached and equilibrated, the
231 experiment is started by upward movement of the sample holder using the outer magnet.
232 In this way, the capsule is — within seconds — brought into the top part of the TZM
233 autoclave, which is in the hot zone of the furnace. Quenching is done the opposite way by
234 sliding the magnet down to bring the capsule back into the water-cooled zone of the
235 assembly, resulting in an almost immediate quench. During the experiments,
236 temperature is monitored using the internal thermocouple of the furnace and an external
237 K-type thermocouple which is inserted into a small bore in the sheath of the TZM
238 assembly and touching the top of the internal autoclave, which is the position closest to
239 the capsule. The pressure is monitored with a factory-calibrated Bourdon-tube gauge and
240 a digital pressure transducer. Considering the small thermal volume of the capsule
241 relative to the massive TZM autoclave we expect no significant thermal gradient within
242 the diffusion couple capsule.

243 Upon the start of the experiment while inserting the rod with the sample capsule into the
244 hot zone of the autoclave, a pressure drop of up to 20 MPa over ca. 10 s is commonly
245 observed. We interpret this to occur due to the displacement and concomitant cooling of
246 a large amount of the hot gas from the autoclave to the cooler parts of the pressure system
247 during to insertion of the rod. The rod is cold relative to the hot atmosphere in the
248 autoclave and upon heating causes the pressure to increase again slowly by typically
249 about 10 MPa over the duration of ca. 50 s. The opposite process is observed upon quench
250 of the experiment, resulting in an up to 20 MPa overpressure during quench, which
251 decreases to ca. 10 MPa overpressure in about 50 s at the end of an experiment. These
252 processes are accounted for by slight over-pressurization before the start of the

253 experiment. The pressure increase at the end of the experiment also ensures that bubbles
254 do not form during the quench. Additionally, it was shown that halogen diffusion is mostly
255 insensitive to pressure variations (e.g., Baker and Balcone-Boissard, 2009). Therefore, we
256 consider these short-lived pressure fluctuations to be insignificant for the results of the
257 diffusion experiments.

258

259 **2.2 Analytical methods**

260 Before preparation for further analyses all samples were weighed individually to ensure
261 no weight loss or gain occurred during the experiment. Diffusion couple capsules were
262 embedded in epoxy and cut open along the center of the capsule and parallel to the
263 diffusion direction (i.e., parallel to the cylindrical capsules rotation axis) using a *Buehler*
264 IsoMet precision saw. All samples were polished using a series of diamond embedded
265 disks to an ultimate grain size of 1 μm while taking care that the polished surface
266 remained parallel to the diffusion direction.

267

268 *2.2.1 Electron probe micro analysis (EPMA)*

269 Major element and halogen concentration analyses of experimental glasses were
270 conducted at the Department of Geoscience of the University of Mainz, and using a *JEOL*
271 JXA8200 electron microprobe. Analyses employed an acceleration voltage of 15 kV, a
272 beam current of 12 nA, and a beam diameter of 10 μm . Dwell times for F and Cl were 120
273 s and 30 s, respectively, yielding a detection limit of 60 ppm (1σ). Analysis dwell times
274 of each major element was: Si 25 s, Al 40 s, Na 20 s, K 30 s, Ca 30 s, Fe 60 s, Mg 30 s, Mn
275 50 s, and Ti 30 s. The device was calibrated before a measurement using the following
276 reference materials: SrF₂ (F), tugtupite (Ca, Na), VG-2 (Ca, Mg), VG-A99 (Fe, Si), MnTi (Mn,
277 Ti), and orthoclase (Al, K). Proper calibration was assessed by analyzing standards VG-2,

278 VG-A99, VG-568 and a natural obsidian standard (~74 wt.% SiO₂) repeatedly during each
279 analytical session. The data was corrected using the ZAF method.
280 Concentration vs. distance profiles were acquired using line-scans with a fixed step-width
281 of 5–250 μm depending on the sample (i.e., the anticipated diffusive speeds and therefore
282 profile distances) and the position of the analyzed points relative to the diffusion
283 interface. Points in the transition zone between the two diffusion couple halves were
284 analyzed using a smaller step-width while the profile ends were covered by a larger step-
285 width. This way it was possible to achieve a high spatial resolution of data, in turn
286 facilitating good resolution of slow diffusion patterns while covering a large range of
287 diffusion widths and keeping analysis time low. Profiles were acquired along traverses
288 away from the capsule edges to avoid analyzing inhomogeneities that could distort
289 diffusion profile data (e.g., devitrification, Feisel et al. 2019).

290

291 2.2.2 Secondary Ion Mass Spectrometry (SIMS)

292 As the concentration of Br and I in the samples is too low to be analyzed by EPMA,
293 experimental samples were further analyzed on the SIMS. The analyses of all HX samples
294 and of CCX1 and CCX4 were carried out during two analytical sessions in October 2018
295 and January 2019 at the University of Lausanne, Switzerland. The other samples (CCX3,
296 CCX5 and all CCX-H) were analyzed at Heidelberg University, Germany, in two sessions
297 in May 2021 and April 2022. Both facilities feature a CAMECA IMS 1280HR ion
298 microprobe. The analytical parameters for all SIMS analyses (halogens and H₂O) in both
299 labs are summarized in Table S1 of the supplementary material.

300

301 Analyses carried out at the University of Lausanne were referenced internally on each
302 sample in an area far outside of the region of diffusion where the constant Si-content was

303 known from EPMA analyses. These data included the less abundant Si-isotope ^{30}Si for
304 calculation of ratios with Br and I which are of low concentration in our samples
305 compared to F and Cl. ^{28}Si and ^{19}F intensities were detected using a Faraday-Cup. The
306 other halogens (^{35}Cl , $^{79/81}\text{Br}$, ^{127}I) and ^{30}Si were analyzed using electron multipliers.
307 The SIMS analyses of halogens in Heidelberg were calibrated using an obsidian glass (
308 JV1, Pichavant, 1987) for F and Cl and the GSE-1G glass for Br and I (Marks et al., 2017a;
309 Marks et al., 2017b; Wiedenbeck, 2017). The accuracy of these analyses is limited by the
310 poor homogeneity of halogens in GSE-1G and the fact that JV1 is not an established
311 reference material with multiple independent determinations of the concentrations. The
312 halogen concentrations determined by SIMS must therefore be interpreted as semi-
313 quantitative with an estimated accuracy of $\leq 50\%$ (relative error).

314

315 For each water-bearing sample (CCX-H), H_2O concentration profiles were acquired next
316 to the halogen profiles using a *CAMECA* IMS3f ion microprobe at Heidelberg University.
317 The H_2O background caused by in-situ contamination was reduced by using a LN_2 -cooled
318 metal plate in the sample chamber. The apparent H_2O concentration caused by in-situ
319 contamination was monitored by analyzing San Carlos olivine and was $< 0.01\text{ wt.}\%$. H_2O
320 analyses were calibrated using JV1 as reference material. Accuracy is estimated to be $<$
321 20% relative error.

322

323 **3. Results**

324 **3.1 Chemical composition of starting materials**

325 The major element compositions of the different starting materials as determined by
326 microprobe analyses are compiled in Table 1. The synthesized samples reveal a similar
327 composition, both plotting in the field of Trachyte in a TAS diagram, very close to the

328 borders to rhyolite and dacite. However, the CCX melt is slightly enriched in alkalis
329 (Na_2O , K_2O) compared to the HX melt. All samples were analyzed using line scans on the
330 microprobe not only to measure F and Cl gradients but also to validate major element
331 homogeneity over the diffusion couple (Fig. 1). Maximum absolute concentrations of F
332 are around 0.5–0.6 wt.% in the HX samples, and about 0.8–1 wt.% in the CCX samples. Cl
333 concentrations are about 0.2–0.3 wt.% in the HX samples and about 0.3–0.4 wt.% in the
334 CCX samples. Absolute Br and I concentrations were only measured for the CCX samples
335 using the SIMS at Heidelberg University. The maximum Br and I concentrations in the
336 halogen-enriched glasses of the CC melts are ~2500 ppm and ~300 ppm, respectively.

337

338 **3.2 Post-experimental diffusion couple textures**

339 During data acquisition, all diffusion couples were assessed for textural homogeneity by
340 means of backscattered electron images. Most samples comprised a clean and straight
341 diffusion interface between two homogeneous glassy areas which were particularly
342 prominent in the anhydrous samples. Some samples and mainly those of the hydrous
343 series contained minor cracks in the glass, which were interpreted to stem from the
344 relatively fast quenching rates (Fig. 2a). However, the cracks were not found to influence
345 the acquired diffusion data. Few samples underwent deformation during the experiments
346 caused by bubble formation or the onset of convection. Both phenomena were mostly
347 recognized in experiments of the water-bearing series in which the melt viscosity was
348 reduced due to the presence of water (e.g., Giordano et al. 2008) which enhanced
349 convection. These effects resulted in distorted diffusion couple geometry, and according
350 samples were discarded from further analysis. Small amounts of oxide phases ($<100\mu\text{m}$;
351 $<<1$ vol.%) were recognized in the halogen-bearing parts of some experiments utilizing
352 the CCX melt. These were analyzed qualitatively using EDS which showed they are

353 tungsten-oxides. This was interpreted to stem from a slight contamination of the sample
354 powder caused by using a tungsten-carbide mill for sample preparation. However, these
355 oxides are not thought to affect halogen diffusion in our experiments.

356

357 **3.3 Halogen diffusion coefficients**

358 Most concentration profiles comprise a smooth and symmetric transition between the
359 high and low halogen concentration portions of the charge, indicating that no other
360 processes than diffusion (e.g., convection) were involved (Fig. 1). Smooth and symmetric
361 profiles are recognized especially for experiments run at high temperatures. Diffusion
362 profiles of Br and I are in general shorter than those of F and Cl. However, all
363 concentration vs. distance profiles presented here yielded good results during fitting and
364 are treated as effective binary diffusion of each halogen element in the melt matrix (e.g.,
365 Zhang 2010).

366 Diffusion coefficients for each acquired concentration vs. distance profile were calculated
367 by fitting the respective data to the equation for constant one-dimensional diffusivity
368 between two semi-infinite media (Crank, 1975):

$$C(x, t) = \frac{C_{low} + C_{high}}{2} + \frac{C_{low} - C_{high}}{2} \operatorname{erf}\left(\frac{x - x_0}{2\sqrt{D t}}\right) \quad (1)$$

369 The parameter $C(x, t)$ describes the concentration at distance x (m) after the experimental
370 time t (s). C_{low} (ppm) is the concentration of the halogen-poor, and C_{high} (ppm) the
371 concentration of the halogen-rich half-space. The term erf indicates the error function. x_0
372 (m) describes the position of the diffusion interface and is in this case determined
373 mathematically. The term D ($\text{m}^2 \text{s}^{-1}$) is the diffusion coefficient.

374 Data fitting was performed using the nonlinear least-squares method implemented in a
375 *MatLab* script, specifically programmed for easy use with diffusion profile data. The

376 results obtained by this script were regularly validated by comparing to the results
377 calculated by the commercially available curve-fitting software *CurveExpertPro*. As
378 equation (1) is independent of absolute concentrations it was used with both,
379 concentration data determined by EPMA and SIMS and ratios of halogen- over silica-
380 concentration (e.g., F/Si) as determined by SIMS to calculate diffusion coefficients. All
381 parameters of equation (1) except x and t were calculated based on the fitting algorithm.
382 In some cases, parameters C_{low} and C_{high} were set manually to facilitate proper fitting
383 calculations, for example where the analyzed profiles did not cover the whole length of
384 diffusion for that particular profile. Additionally, some profiles suffered from slight
385 distortion, caused, for example, by deformation of the capsule during quench, and
386 resulting in asymmetrical profiles which was mainly prominent in the longest profiles of
387 F diffusion. Where profiles were asymmetrical, each half of the diffusion profile was fitted
388 individually using equation (1) and the results were compared to those of the complete
389 profile. Those profiles of which the individually fitted halves yielded different diffusion
390 coefficients larger than 0.6 orders of magnitude were discarded. F diffusion coefficients
391 calculated on profiles where this method was applied are indicated by "*" in Table 2.
392 Typical concentration vs. distance profiles of sample CCH-X-42 fitted with equation (1)
393 are illustrated in Figure 2b. All measured diffusion profiles with their respective fit curves
394 are provided in Figures S1-S4 of the supplementary material.

395

396 3.3.1 Anhydrous HX

397 Of the nine HX experiments presented in Feisel et al. (2019), six samples were analyzed
398 by SIMS for the scope of this study. One concentration vs. distance profile was acquired
399 for each of the samples, covering lengths of ~500–1000 μm . Diffusivities of F and Cl
400 determined here agree well with the results of microprobe analyses indicating that F is

401 generally faster than Cl by up to 2 orders of magnitude in the investigated temperature
402 range. D values are on the order of 10^{-15} to 10^{-13} $\text{m}^2 \text{s}^{-1}$ for fluorine and 10^{-17} to 10^{-15} $\text{m}^2 \text{s}^{-1}$
403 for chlorine, of which the lower and higher values each correspond to temperatures of
404 750 °C (HX9) and 950 °C (HX8), respectively (Table 2). Bromine and iodine diffusivities
405 are generally lower than those of chlorine, with Br consistently being faster than I. D_{Br}
406 ranges from 10^{-18} to 10^{-15} $\text{m}^2 \text{s}^{-1}$ while D_{I} values are about half an order of magnitude
407 smaller at the corresponding temperatures, ranging from ca. 10^{-18} to 10^{-16} $\text{m}^2 \text{s}^{-1}$ (Table
408 2).

409

410 3.3.2 Anhydrous CCX

411 Out of seven anhydrous experiments performed using CCX melt and analyzed by EPMA,
412 three were discarded due to textural features indicating compromised geometry. The
413 remaining four experiments were additionally analyzed by SIMS. Diffusivities of all
414 halogens in samples using the CCX melt are similar to those obtained for the HX melt.
415 However, the difference between F and Cl diffusion is less pronounced in this melt
416 composition compared to HX. Diffusion coefficients of F (D_{F}) range from $\sim 10^{-14}$ to 10^{-13}
417 $\text{m}^2 \text{s}^{-1}$ in the temperature range of 850 °C (CCX3) to 1000 °C (CCX5). Cl diffusion is up to
418 two orders of magnitude slower with $\sim 10^{-17}$ to 10^{-14} $\text{m}^2 \text{s}^{-1}$ at these temperatures. For
419 these two halogens, results of EPMA and SIMS analyses agree well within statistical errors
420 (Table 2). Diffusion of Br is slower than that of Cl by ~ 0.5 – 1.5 orders of magnitude and I
421 is even slower than Br by ~ 1 order of magnitude with diffusivities in the range of $\sim 10^{-16}$ –
422 10^{-14} and $\sim 10^{-16}$ – 10^{-15} $\text{m}^2 \text{s}^{-1}$, respectively.

423

424 3.3.3 Hydrous CCX-H

425 All diffusion couple experiments of the hydrous series performed in the TZM assembly
426 were analyzed both by EPMA (F, Cl) and by SIMS (F, Cl, Br, I). In general, diffusivities are
427 higher than those obtained for the anhydrous CCX series with D_F ranging from $\sim 10^{-12}$ to
428 $\sim 10^{-11} \text{ m}^2 \text{ s}^{-1}$ from 1000–1200 °C. Cl diffusion is 0.5 to 1 orders of magnitude slower than
429 F, and Br and I are even slower. However, diffusivities of Cl, Br and I are all within less
430 than one order of magnitude of each other at the respective experimental temperatures
431 (Table 2). This results in a total diffusivity range of only 1–2 orders of magnitude among
432 all halogens which is significantly less than observed for the anhydrous series of CCX melt
433 which spans a range of 3–4 orders of magnitude among all halogens (Table 2).

434

435 **3.4 Temperature dependence of halogen diffusion**

436 All diffusion coefficients determined in this study increase with increasing temperature
437 for each of the used melt compositions. As shown in various previous studies, diffusivity
438 follows Arrhenian behavior (e.g., Baker and Balcone-Boissard 2009), which is also
439 confirmed by our data. Plotted in a diagram of $\log(D)$ vs. inverse temperature (Fig. 3), the
440 data clearly shows this Arrhenian trend and can be described by the following equation:

$$D = D_0 * e^{-\frac{E_A}{R*T}} \quad (2)$$

441 where D is the diffusion coefficient ($\text{m}^2 \text{ s}^{-1}$), D_0 is the pre-exponential factor, E_A the
442 activation energy (J/mol), R the universal gas constant ($8.3145 \text{ J mol}^{-1} \text{ K}^{-1}$), and T is the
443 temperature (K). The characteristic parameters describing the diffusive behavior of each
444 halogen in a specific melt composition are D_0 and E_A which are mathematically
445 represented by the y-axis intersect and the slope of the trendline in the Arrhenius
446 diagram, respectively. For the calculation of the characteristic Arrhenius parameters,
447 equation (2) was linearized to the form

$$\log(D) = -\frac{E_A}{\ln(10) * R * T} + \log(D_0) \quad (3)$$

448 and fitted to all available data (EPMA + SIMS) of each starting material using the nonlinear
449 least-squares method implemented with *Matlab*. Where multiple profiles were measured
450 on one sample and with the same method, the resulting diffusion coefficients were
451 averaged to avoid a weighing effect during the Arrhenius fitting calculations. The results
452 of all fits are illustrated in Fig. 3 and listed in Table 3. Fitted with the data of the two
453 anhydrous series, the results of the HX and CCX starting materials are similar, yielding
454 activation energies in the range of ~200–290 kJ/mol for all halogens. However, diffusion
455 in HX melt generally comprises slightly higher activation energies than in CCX melt. The
456 similarity is further represented by the parallel fit curves in the Arrhenius diagrams (Fig.
457 3). However, in the HX series it is apparent that F is significantly faster than all other
458 halogens which is represented by a gap of 1.5–2 orders of magnitude between the fit lines
459 of F and Cl, while Cl, Br and I are all within about 1.5 orders of magnitude. In comparison,
460 the data of the CCX series shows more equally spaced Arrhenius fits of the halogens
461 spanning an overall similar range of diffusivities compared to the HX experiments.
462 The data of the hydrous series shows the same ordering of diffusive speeds, with F being
463 the fastest and I being the slowest, but spanning a narrower range of diffusivities among
464 all halogens. In the Arrhenius diagram this is represented by the more closely spaced fit
465 curves compared to the anhydrous data. Activation energies are similar to the anhydrous
466 series, except for F whose E_A is only 131 kJ/mol in the hydrous case compared to 198–
467 213 kJ/mol in the anhydrous experiments (Table 3).

468

469 **3.5 H₂O content in CCX-H experiments**

470 The experimental products of the CCX-H series show a range of 1.5–2 wt.% H₂O (Table
471 2). Most of the hydrous experiments suffered from water loss to some extent, which was
472 indicated by increasing totals towards one end of the diffusion profile recognized in the
473 EPMA results (Fig. 1) and confirmed by SIMS analyses (Fig. 2; Fig. S1 of the
474 supplementary material). The use of CH₄ in the pressure medium helped to reduce this
475 issue but did not resolve it completely. Interestingly, water was only lost on the halogen-
476 depleted side of the diffusion couple leading to a weak, mostly linear H₂O-concentration
477 gradient over the length of the diffusion couple. Experiments comprising significant
478 water loss were discarded, however, small losses were considered negligible with regard
479 to diffusivity. This is particularly true considering that the transition between high and
480 low halogen concentration in most cases is relatively narrow, especially relative to the
481 short diffusion lengths of the slower diffusing halogens. Over this narrow range the H₂O
482 concentration can be considered to be constant within error. Due to this issue, we have
483 calculated the average H₂O-concentration over the length of the transition zone of each
484 hydrous diffusion profile and reported it with other relevant data in Table 2.

485

486

487 **4. Discussion**

488 **4.1 Halogen diffusion mechanisms**

489 Diffusivity and the mechanism of halogen diffusion are both strongly dependent on the
490 composition and internal structure of the host glass. In the case of aluminosilicate melts
491 a vast range of compositions are present in nature and structural melt parameters such
492 as the degree of polymerization show an accordingly wide range. Arrhenius parameters
493 of halogen diffusion have been determined for a range of different melt compositions and
494 illustrate the strong compositional dependence of halogen diffusion (Watson and Bender,

495 1980; Dingwell and Scarfe, 1984; Dingwell and Scarfe, 1985; Bai and Koster van Groos,
496 1994; Alletti et al., 2007; Balcone-Boissard et al., 2009; Böhm and Schmidt, 2013; Fortin
497 et al., 2017; Yoshimura, 2018; Balcone-Boissard et al., 2020). Most of these studies
498 studied F and Cl diffusion and only few have investigated Br diffusion. However, none of
499 the available studies provide I diffusivity data. Hence the present study is the first to our
500 knowledge.

501 One major compositional parameter defining the behavior of silicate melts is the degree
502 of polymerization, which itself is a function of, for example, the amounts of SiO₂, volatiles
503 or network modifying cations. Typically, the degree of polymerization in a melt is
504 characterized by the parameter of NBO/T, i.e., the ratio of non-bridging oxygen atoms
505 over tetrahedrally coordinated ions (Mysen, 1988). The melts of the present study exhibit
506 NBO/T values in the range of ~ 0.08–0.11, calculated using the method of (Mysen, 1988)
507 and splitting FeO_{tot} equally between Fe₂O₃ and FeO (Giordano et al., 2008). The NBO/T
508 values are reported together with the bulk rock chemistry in Table 1.

509

510 The bond strength of an ionic species dissolved in melt generally increases with the
511 charge of that species (Zhang et al., 2010). Halogens in the combined state (e.g., as a
512 halide) typically have an univalent charge of -1 and are therefore comparably weakly
513 bonded to other atoms, despite being very reactive. This means that halogens are
514 relatively easily detached from their original site by breakage of the ionic bonds around
515 a previously coordinated halogen ion before moving to a new site (Zhang et al., 2010).
516 Additionally, even though the halogen group elements have the highest electronegativity
517 of all elements in their respective periods, the significant within group variation in this
518 parameter means that the bond strength should also vary accordingly. For example, F,
519 being the most electronegative, should foster the highest bond strength of the halogens.

520 Within this simplified chemical framework the diffusive “jump” can be envisioned to
521 proceed by passing through an aperture formed by neighboring ions. The size and
522 stiffness of the aperture should generally be structurally controlled and relate to the
523 polymerization of the melt, while the detachment rate should be controlled by ionic
524 charge (Zhang et al., 2010). The size and number of the apertures are determined by the
525 ions present in the melt and the resulting “free space” formed which is called the ionic
526 porosity. Naturally, smaller ions such as F diffuse faster, because they can pass through
527 apertures of various sizes and are therefore less sensitive to ionic porosity. Larger ions
528 can only pass large apertures and their diffusion-rate is therefore structurally controlled
529 and, hence, more sensitive to ionic porosity. This is typically the case for noble gases and
530 monovalent ions (e.g., Henderson et al., 1985; Lux, 1987; Alletti et al., 2007) and was also
531 confirmed for halogens by recent studies which showed that diffusion is negatively
532 correlated with the ionic radius of the halogen in low NBO/T melts, i.e., in melts with a
533 high degree of polymerization (Rhyodacite/Trachyte, NBO/T: ~ 0.08 , Feisel et al., 2019;
534 Na- and K-Phonolite, NBO/T: ~ 0.08 , Balcone-Boissard et al., 2020). In melts with higher
535 NBO/T (lower degree of polymerization) halogen diffusion typically comprises only weak
536 correlation with the ionic radius. Alletti et al. (2007) suggested that halogen diffusion in
537 a basaltic melt (NBO/T ~ 0.5) is mostly independent of “free space” in the melt structure
538 but instead occurs by exchange mechanisms that are predominantly controlled by the
539 preservation of local charge balance, similar to the diffusion of divalent and trivalent
540 cations (Henderson et al., 1985). This means that in this kind of melt movement of
541 halogens is mainly rate-limited by the charge-controlled detachment process.

542

543 The present study complements the data of Feisel et al. (2019) by including all new SIMS
544 analyses which also cover Br and I in two different highly polymerized melts (HX: NBO/T

545 ~0.08; CCX: NBO/T ~0.08–0.11). Figure 4 illustrates the diffusivity of the halogens in
546 different melt compositions relative to their ionic radii (Shannon, 1976) at 1250 °C. The
547 data of Alletti et al. (2007) are included as a reference of a weakly polymerized melt and
548 the data of studies performed at lower temperatures were calculated based on the
549 published Arrhenius parameters. The results of the present study and especially those of
550 Br and I confirm the strong influence of the ionic radius on diffusivity in polymerized
551 anhydrous melts. The difference between F and Cl is particularly pronounced for most
552 anhydrous studies, yielding a significant negative slope between the two in the diagram
553 (Fig. 4). However, the slope between Cl and Br defines a discordance to the trend between
554 F and Cl in most studies. For the anhydrous CCX melt the slope increases from Cl to Br
555 and this trend is continued to I. For the anhydrous HX melt, however, the slope decreases
556 from Cl to Br and I. This stepwise correlation of ionic radius and diffusivity is recognized
557 for all polymerized melts that were analyzed for Br (and I). We interpret this behavior to
558 show that the ionic porosity of the melt is limited to distinct aperture sizes dictated by
559 the melt structure. The fact that Cl and Br seem to have a similar diffusivity in anhydrous
560 phonolite (Balcone-Boissard et al., 2009) regardless of their different ionic radii suggests
561 that Cl and Br move through similarly sized apertures in this melt. By contrast, in the HX
562 and especially CCX melt, Br cannot move through the same apertures as Cl due to its size,
563 resulting in its lower diffusivity. Likewise, the I ion is too large to pass these aperture
564 sizes and can diffuse only in the vicinity of ions that allow the formation of larger openings
565 between sites. F, being the fastest diffuser among all of the halogens, can move through
566 both small openings and all the larger apertures that are used by Cl, Br, and I in a
567 polymerized anhydrous melt.

568

569 **4.2 The effect of H₂O on diffusion**

570 The hydrous experiments (~ 1.5 wt.% H_2O) generally yield high diffusivities which span
571 a range of only 1–2 orders of magnitude among the different halogens (Fig. 3). This
572 restricted range of diffusivities contrasts those of the anhydrous CCX melt which displays
573 3–4 orders of magnitude variation. For better comparison to the anhydrous results, the
574 Arrhenius parameters of the anhydrous CCX series were used to calculate hypothetical
575 anhydrous diffusivities in the temperature range of the hydrous experiments and plotted
576 in an Arrhenius diagram together with the hydrous Arrhenius fits (Fig. 5). The calculated
577 anhydrous values are consistently lower than the hydrous data. While the difference for
578 F is only ~ 0.2 log units at high T (1200 °C) and ~ 0.8 log units at low T (1000 °C), the
579 difference for the slower diffusing halogens is more pronounced. Cl is ~ 1 log unit, Br
580 ~ 0.5 – 1.5 log units and I ~ 2 log units faster in hydrous compared to dry melt. These
581 results highlight the strong influence of H_2O on halogen diffusivity in the high silica melts
582 of this study. These findings are in contrast to the those of Balcone-Boissard et al., (2020),
583 who showed that the addition of up to 2.65 wt.% H_2O exerts only minor influence on Br
584 diffusion in phonolitic melts; a fact which further underlines the strong compositional
585 dependence of halogen diffusion.

586

587 The increase in diffusivity in the melts of this study indicates a change in the diffusion
588 mechanism that is likely caused by the modification of the melt structure due to the
589 addition of H_2O . This effect can also be observed in Figure 4, where the strong correlation
590 of diffusivity with the ionic radii observed in the anhydrous melts is highly attenuated for
591 the hydrous case. Indeed, the hydrous data is more similar to those of basalt confirming
592 that water effectively lowers the degree of polymerization of the melt. In doing so, the
593 rate-limiting influence of melt structural apertures is reduced which makes the process
594 of detachment more relevant as a diffusion-rate-limiting factor. Interestingly, the slope

595 between Cl and Br increases in the hydrous melt similar to the behavior in the anhydrous
596 CCX melt. At the relatively low water contents of ~1.5 wt.% used in the present study, the
597 majority of the added H₂O is expected to be dissociated to form hydroxyl (OH) in the melt
598 (e.g., Silver et al., 1990; Ihinger et al., 1999). The experimental results indirectly show that
599 hydroxyl effectively depolymerizes the melt and allows for faster halogen diffusion.
600 However, even in the hydrated melt, halogen diffusion is still somewhat correlated with
601 the ionic radii. This may indicate that for the given added water content (~1.5 wt.%) the
602 melt structure is still polymerized enough to assert some control on the aperture-limited
603 jumping processes, mainly affecting larger ions. We speculate that diffusivities will
604 increase even more and may become more equal among the halogen species with
605 increased water content. Indeed, our own preliminary results using 3.5 wt.% of H₂O
606 added to the same CCX melt and in the same experimental setup (e.g., 1000 °C / 160 MPa)
607 yield still higher diffusivities that span over an even smaller range among all the halogens
608 (Fig. 5). This strongly implies that the rate at which a halogen ion jumps into a new site
609 grows as the melt becomes increasingly depolymerized.

610

611 Another notable distinction between anhydrous and hydrous melt is the change in the
612 activation energy, E_A , of F diffusion, which is lowered from 198 to 131 kJ mol⁻¹, while the
613 E_A of the other halogens are only weakly affected (Table 3). That is, F diffusion becomes
614 less sensitive to temperature in hydrous melts, which is apparent in the comparably
615 shallow slope of the F Arrhenius fit in Figures 3c and 5. As already concluded from the
616 observation that F diffusion is less strongly enhanced in the hydrous melt, compared to
617 the other halogens, F diffusion appears only weakly sensitive to the melt structure and is
618 more strongly controlled by the detachment process. F is known to substitute for bridging
619 oxygen in Si-O-Si units (e.g., Mysen and Virgo, 1985) and preferably coordinates with Si

620 or Al (e.g., Zeng and Stebbins, 2000; Liu and Nekvasil, 2002) , while Cl and Br prefer to
621 coordinate in larger sites (Cassidy et al., *in press*). Similarly, OH⁻ typically replaces oxygen
622 bonds of the Si-O network (Mysen et al., 1980). Consequently, the addition of H₂O and the
623 concomitant formation of OH⁻ in the melt must influence the preferred bonding
624 environment of F in the silicate network. Here, we interpret the change in activation
625 energy of F to reflect that the rate of detachment from its site is enhanced by the
626 replacement of F with OH⁻. This way more F is available for diffusion compared to the
627 anhydrous case, which in turn suggests that F diffusion becomes less sensitive to
628 temperature (i.e., lower activation energy) under hydrous conditions. Consequently, we
629 speculate that the introduction of water has a two-fold effect on the halogen diffusion
630 mechanism: 1) The resulting decrease in the degree of polymerization generally
631 promotes halogen diffusivities by increasing the ionic porosity; 2) The strong effect on
632 the activation energy of F is interpreted to be due to OH⁻ directly affecting the bonding
633 environment of F. Clearly, these explanations are qualitative, and therefore meant to
634 suggest possible microscale chemical processes that might explain our macroscopic
635 observations. Testing these ideas could benefit from appropriate imaging experiments
636 (e.g., Schaller et al., 1992; Zeng and Stebbins, 2000; Mysen et al., 2004) that may isolate
637 the relevant coordination and bonding environments of the halogens, with implications
638 for their bond strengths in natural silicate melt.

639

640 **4.3 Halogen diffusion and viscosity**

641 Oxygen self-diffusion is commonly seen as a solid estimate of the structurally defined
642 lower limit of diffusion rates in silicate melts (Oishi et al., 1975; Shimizu and Kushiro,
643 1984; Dingwell, 1990; Leshner, 2010; Zhang and Ni, 2010) and appears closely related to
644 anhydrous melt viscosity via the Eyring equation (Glasstone et al., 1941; Dingwell, 1990;

645 Zhang and Ni, 2010). Calculated oxygen self-diffusivity of our anhydrous samples using
646 viscosity data derived from the model of Giordano et al. (2008) are included in Figure 3.
647 Diffusion of all four halogens in the dry melt is at least 1 order of magnitude faster than
648 the calculated oxygen self-diffusion (Fig. 3), indicating that the silicate network behaves
649 as in quasi-static structural sites relative to the diffusing halogen ion; therefore, diffusion
650 may only be affected weakly by self-diffusion of the silicate network. Notably, iodine
651 diffusion is the slowest of the halogens and thus closest to the calculated Eyring diffusion.
652 The steeper slope of the Eyring diffusivity compared to the slope of the halogen fits in the
653 Arrhenius diagrams (Fig. 3) suggests that with increasing temperature the jump
654 frequency of iodine may approach that of the Si-O bonds in the silicate structure and the
655 network surrounding iodine may no longer be viewed as quasi-static. With increasing
656 temperature, iodine diffusion might therefore be enhanced by the cooperative
657 rearrangement of the local silicate structure as oxygen diffuses.

658 The activation energy of diffusion, i.e., the slopes of the Arrhenius fits of the other
659 halogens (Fig. 3) likewise indicate that all fits will meet the lower limit of oxygen self-
660 diffusion at higher temperatures. Consequently, at sufficiently high temperatures (i.e.,
661 sufficiently low viscosity), all halogens will move with the same frequency as the silicate
662 network, hence, diffusion will be the same for all halogens, being determined by the jump
663 frequency of the structural movement of the Si-O bonds, broadly consistent with the
664 conclusions of Dingwell (1990).

665 The indicated oxygen self-diffusivities in Figure 3 are considered to represent the lower
666 limit of possible values as they are based on the halogen-depleted melt compositions. In
667 F bearing melts, like those of the enriched diffusion couple halves, the viscosity can be
668 significantly reduced due to the depolymerizing effect of F (e.g., Dingwell et al., 1985;
669 Giordano et al., 2008), which will in turn yield higher calculated oxygen self-diffusivities.

670 For the melts used in this study, the F-bearing melts yield Eyring diffusivities of up to 1
671 order of magnitude higher than the F-depleted melts, approaching the derived
672 diffusivities of iodine at the highest of the experimental temperatures. Therefore, during
673 the experiments, the melt structural lower limit of halogen diffusion is expected to change
674 upon diffusion of F. In the resulting F enriched melts, the halogen diffusivities might
675 approach the Si-O jump frequency already at lower temperatures compared to the F-
676 depleted case. Expanding on this idea, we speculate that when F diffusivity and the F-
677 depleted Eyring diffusivity are within ~1 order of magnitude in Arrhenius space (i.e., at
678 sufficiently high temperatures), diffusion of F into the depleted melt will enhance F
679 diffusivity by decreasing the degree of polymerization. This would in turn result in a
680 concentration dependent F diffusivity comparable to the case of H₂O (e.g., Behrens et al.,
681 2004). However, it is unclear if the Eyring equation is valid for the case of F-enriched
682 polymerized melts.

683

684 In case of the hydrous melt, the Eyring equation was shown to be inapplicable as it under-
685 estimates oxygen-diffusion by 3–4 orders of magnitude in rhyolitic melt (Behrens et al.,
686 2007). In this case, oxygen no longer moves solely by self-diffusion but is mainly
687 transported by the diffusion of molecular H₂O in the melt (Behrens et al., 2007; Zhang and
688 Ni, 2010) which shows that the structural limit of diffusion rates in hydrous melt can no
689 longer be described by the self-diffusion of oxygen.

690 In an attempt to define a lower limit of diffusion rates in the hydrous melt of the present
691 study, calculated oxygen self-diffusivities of rhyolite and dacite (Tables 4 and 5 of Zhang
692 and Ni, (2010)) have been included in Figure 3c. However, these predicted self-
693 diffusivities are faster than the experimentally constrained diffusivities of Br and I over a
694 large range of the investigated temperatures. This is interpreted to indicate that in a

695 slightly depolymerized melt of sufficiently large ionic porosity, the jumps of Si-O bonds
696 no longer define the lower limit of transport rates but that the diffusive characteristics
697 may be mostly determined by the ionic charge and radius.

698

699 **5. Conclusions**

700 The diffusion characteristics of four halogens in silicic melts were determined
701 experimentally under anhydrous and hydrous conditions and over a temperature range
702 relevant for magmatic processes (750–1200 °C). The results are consistent in that F is the
703 fastest diffusing halogen and diffusivity decreases with increasing ionic radius indicating
704 that diffusivity is highly dependent on the melt structure and therefore melt composition.
705 Halogen diffusion covers a range of 3–4 orders of magnitude in silica-rich anhydrous
706 melt, but diffusivities increase dramatically when water is added. The diffusivity increase
707 due to the addition of water is most pronounced for the largest of the halogens and thus,
708 the slowest diffusing halogen iodine. The effect of added water is weaker for the smaller
709 halogens, leading to an overall diffusive range of only 1–2 orders of magnitude among all
710 halogens. While the temperature dependence in terms of activation energies of Cl, Br and
711 I diffusion is similar in the dry and wet case ($\sim 200\text{--}290$ kJ/mol), E_A of F is reduced by
712 about 35% in hydrous (~ 131 kJ/mol) compared to anhydrous melt (~ 200 kJ/mol). This
713 is interpreted to indicate that F diffusion is more strongly dependent on the ionic
714 detachment during diffusion and is therefore less sensitive to ionic porosity, especially in
715 hydrous melt.

716 This study represents a significant contribution to the experimental database of halogen
717 diffusion in natural silicate melts, and to our knowledge, is the first to include coherent
718 results on iodine diffusion. The new data are highly relevant for the better understanding
719 of halogen mobility in magmatic processes. Our results suggest that diffusive

720 fractionation of faster diffusing components such as H₂O and the halogens, could occur
721 during bubble formation in a weakly water-enriched magma during slow magma ascent
722 (e.g., Watson, 2017). Modeling of different magma ascent and bubble growth scenarios
723 using the proposed diffusion parameters may help to characterize possible halogen
724 fractionation trends during volcanic unrest.

725

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734

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916 **Tables**

917 **Table 1:** Major element composition of the synthesized starting glasses as determined by EPMA
 918 and SIMS analysis. All values represent units of wt.% unless otherwise stated. All data are
 919 corrected for Na-loss during EPMA analysis.

	HX ^a		CCX	
	enriched (n = 15)	depleted (n = 14)	enriched (n = 6)	depleted (n = 5)
SiO ₂	68.5 (4)	68.3 (3)	67.8 (3)	67.9 (5)
TiO ₂	0.46 (2)	0.46 (2)	0.66 (3)	0.67 (2)
Al ₂ O ₃	14.7 (1)	14.57 (8)	14.49 (4)	14.48 (7)
FeO	4.6 (3)	4.76 (6)	3.84 (5)	3.94 (8)
MgO	0.46 (2)	0.44 (2)	0.60 (3)	0.62 (2)
MnO	0.13 (2)	0.15 (2)	0.12 (4)	0.11 (2)
CaO	2.85 (7)	2.82 (5)	2.11 (3)	2.12 (3)
Na ₂ O	6.2 (2)	6.35 (7)	6.9 (1)	7.8 (1)
K ₂ O	2.10 (3)	2.20 (3)	2.39 (3)	2.65 (4)
F	0.53 (2)	0.04 (3)	0.96 (2)	0.03 (1)
Cl	0.26 (1)	0.02 (2)	0.4 (1)	0.04 (1)
Br	nd	nd	~ 2500 ppm	~ 1 ppm
I	nd	nd	~ 290 ppm	0
Total	100.52	100.05	99.82	100.35
NBO/T ^c	0.08	0.08	0.09	0.11

^a Data from Feisel et al. (2019)

^b Absolute values of Br and I estimated based on SIMS analyses and only available for CCX melts.

^c NBO/T is calculated after (Mysen, 1988) and Giordano et al. (2008).

921 **Table 2:** Conditions and results of all experiments. Experiments of the anhydrous series (HX and CCX) were carried out at $P = 1$ bar. Hydrous experiments
 922 were done at $P = 160$ MPa. All diffusivities in $m^2 s^{-1}$.

sample	T [° C]	t [s]	H ₂ O [wt.%]	#	D_F	s.d.	D_{Cl}	s.d.	D_{Br}	s.d.	D_I	s.d.
HX9	750	3023535	-	1 ^a	1.6×10^{-15}	4.3×10^{-16}	1.7×10^{-17}	8.4×10^{-18}				
				2 ^b	$1.6 \times 10^{-15*}$	3.2×10^{-16}	3.9×10^{-17}	1.1×10^{-17}	3.4×10^{-18}	8.8×10^{-19}	1.8×10^{-18}	7.8×10^{-19}
HX7	800	587235	-	1 ^a	6.6×10^{-15}	1.6×10^{-16}	4.9×10^{-17}	1.4×10^{-17}				
				2 ^a	1.1×10^{-14}	1.9×10^{-15}	1.6×10^{-16}	4.2×10^{-17}				
				3 ^b	1.1×10^{-14}	6.0×10^{-16}	4.0×10^{-16}	6.2×10^{-17}	2.3×10^{-17}	4.7×10^{-18}		
HX3	850	171502	-	1 ^a	1.6×10^{-14}	5.0×10^{-15}	3.9×10^{-16}	2.0×10^{-16}				
				2 ^a	1.9×10^{-14}	5.9×10^{-15}	4.5×10^{-16}	2.4×10^{-16}				
				3 ^b	$2.3 \times 10^{-14*}$	4.6×10^{-15}	4.3×10^{-16}	9.7×10^{-17}	6.2×10^{-17}	2.9×10^{-17}	4.0×10^{-17}	2.1×10^{-17}
HX10	850	266855	-	1 ^a	1.9×10^{-14}	4.4×10^{-15}	3.9×10^{-16}	1.2×10^{-16}				
				2 ^a	2.0×10^{-14}	5.1×10^{-15}	7.4×10^{-16}	4.0×10^{-16}				
				3 ^b	$2.1 \times 10^{-14*}$	1.8×10^{-15}	6.2×10^{-16}	7.4×10^{-17}	8.9×10^{-17}	1.6×10^{-17}	2.7×10^{-17}	3.6×10^{-16}
HX5	900	154560	-	1 ^a	4.7×10^{-14}	6.7×10^{-15}	1.0×10^{-15}	2.0×10^{-16}				
				2 ^a	4.6×10^{-14}	8.9×10^{-15}	8.7×10^{-16}	1.9×10^{-16}				
				3 ^b	$3.2 \times 10^{-14*}$	6.2×10^{-15}	1.0×10^{-15}	1.3×10^{-16}	1.6×10^{-16}	6.5×10^{-17}		
HX8	950	87025	-	1 ^a	1.3×10^{-13}	2.5×10^{-14}	4.5×10^{-15}	1.1×10^{-15}				
				2 ^a	1.1×10^{-13}	1.4×10^{-14}	1.5×10^{-15}	6.7×10^{-16}				
				3 ^b	1.4×10^{-13}	1.6×10^{-14}	5.9×10^{-15}	8.7×10^{-16}	1.3×10^{-15}	2.7×10^{-16}	3.9×10^{-16}	1.9×10^{-16}

^a Results from EPMA analyses. EPMA data of HX series from Feisel et al. (2019)

^b Results from SIMS analyses

*Asymmetric profiles were first fitted each half at a time to confirm reliability of the fit of the complete profile (see text for further discussion)

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sample	T [° C]	t [s]	H ₂ O [wt.%]	#	D_F	s.d.	D_{Cl}	s.d.	D_{Br}	s.d.	D_I	s.d.
CCX3	850	237370	-	1 ^a	$6.6 \times 10^{-14*}$	1.3×10^{-14}	2.4×10^{-15}	9.0×10^{-16}				
				2 ^b	$3.6 \times 10^{-14*}$	1.2×10^{-14}	3.5×10^{-15}	7.1×10^{-16}	4.8×10^{-16}	7.3×10^{-17}		
CCX4	900	170880	-	1 ^a	$6.3 \times 10^{-14*}$	1.2×10^{-14}	1.7×10^{-15}	4.9×10^{-16}				
				2 ^a	$6.6 \times 10^{-14*}$	1.1×10^{-14}	5.7×10^{-15}	1.2×10^{-15}				
				3 ^b	5.4×10^{-14}	1.0×10^{-14}	9.7×10^{-15}	1.4×10^{-15}	1.2×10^{-15}	1.2×10^{-16}	1.7×10^{-16}	5.9×10^{-17}
CCX1	950	82770	-	1 ^a	$1.1 \times 10^{-13*}$	1.6×10^{-14}	8.8×10^{-15}	1.3×10^{-15}				
				2 ^a	$1.2 \times 10^{-13*}$	1.9×10^{-14}	7.1×10^{-15}	1.2×10^{-15}				
				3 ^b	$1.1 \times 10^{-13*}$	1.4×10^{-14}	7.7×10^{-15}	7.5×10^{-16}	1.1×10^{-15}	1.3×10^{-16}	2.5×10^{-16}	1.1×10^{-16}
CCX5	1000	74069	-	1 ^a	6.9×10^{-13}	6.1×10^{-14}	4.0×10^{-14}	5.9×10^{-15}				
				2 ^a	5.0×10^{-13}	4.2×10^{-14}	5.8×10^{-14}	8.2×10^{-15}				
				3 ^a	6.4×10^{-13}	1.4×10^{-13}	6.7×10^{-14}	2.3×10^{-14}				
				4 ^b	6.7×10^{-13}	7.4×10^{-14}	7.0×10^{-14}	2.6×10^{-14}	1.3×10^{-14}	7.5×10^{-15}		
				5 ^b	7.5×10^{-13}	4.4×10^{-14}	7.3×10^{-14}	4.3×10^{-15}	1.2×10^{-14}	1.3×10^{-15}	9.4×10^{-16}	1.3×10^{-16}
CCX-H-38	1000	16380	1.9 (1)	1 ^a	3.3×10^{-12}	4.2×10^{-13}	4.3×10^{-13}	1.4×10^{-13}				
				2 ^b	3.3×10^{-12}	1.6×10^{-13}	5.4×10^{-13}	3.7×10^{-14}	2.5×10^{-13}	1.6×10^{-14}	9.4×10^{-14}	8.1×10^{-15}
CCX-H-40	1050	10200	1.9 (1)	1 ^a	6.2×10^{-12}	9.6×10^{-13}	9.2×10^{-13}	2.2×10^{-13}				
				2 ^b	6.6×10^{-12}	2.6×10^{-13}	1.3×10^{-12}	1.4×10^{-13}	5.8×10^{-13}	6.8×10^{-14}	2.3×10^{-13}	2.9×10^{-14}
CCX-H-42	1100	8100	1.5 (1)	1 ^a	5.5×10^{-12}	7.7×10^{-13}	7.9×10^{-13}	1.6×10^{-13}				
				2 ^b	5.3×10^{-12}	1.3×10^{-13}	9.5×10^{-13}	2.3×10^{-13}	4.2×10^{-13}	3.8×10^{-14}	1.6×10^{-13}	2.4×10^{-14}
CCX-H-43	1150	6000	2.1 (2)	1 ^a	1.3×10^{-11}	1.4×10^{-12}	4.2×10^{-12}	6.7×10^{-13}				
				2 ^b	1.3×10^{-11}	4.4×10^{-13}	3.6×10^{-12}	2.4×10^{-13}	1.6×10^{-12}	9.3×10^{-14}	6.5×10^{-13}	6.7×10^{-14}
CCX-H-44	1200	4800	2.0 (2)	1 ^a	1.9×10^{-11}	2.0×10^{-12}	7.6×10^{-12}	1.3×10^{-12}				
				2 ^b	2.0×10^{-11}	1.1×10^{-12}	7.7×10^{-12}	8.5×10^{-13}	3.9×10^{-12}	4.8×10^{-13}	2.1×10^{-12}	4.2×10^{-13}

929 **Table 3:** Arrhenius parameters of all halogens and oxygen self-diffusion in all investigated melt
 930 compositions. D_0 is given in $m^2 s^{-1}$ and E_A is given in $kJ mol^{-1}$. Calculation of Eyring diffusivities is
 931 described in the text.

	HX			CCX			CCX-H		
	D_0	E_A	+/-	D_0	E_A	+/-	D_0	E_A	+/-
F	1.50e-4	212.7	28	5.11e-5	197.5	99.4	7.76e-7	131.0	38.0
Cl	1.24e-4	245.8	51	5.10e-5	222.6	112	1.74e-4	210.1	67.6
Br	1.78e-3	287.6	66.5	1.16e-5	225.2	367.8	3.97e-5	201.4	153.5
I	3.50e-4	280.0	63.8	2.51e-7	207.4	973.2	1.02e-4	222.2	180.2
O_{Eyring}	6.9e-1	399.1	26.4	1.43e-1	375.5	22.0	-	-	-

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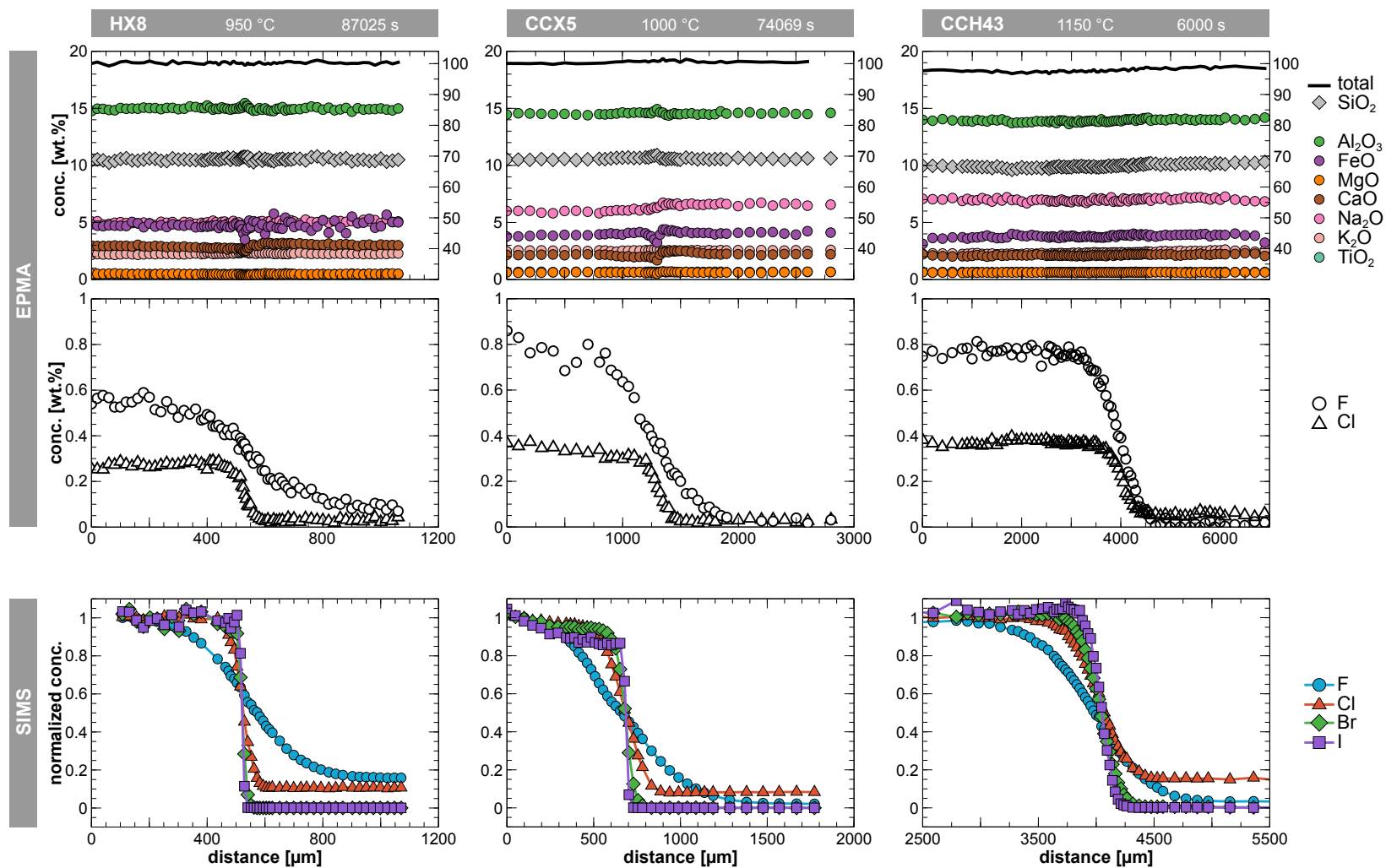
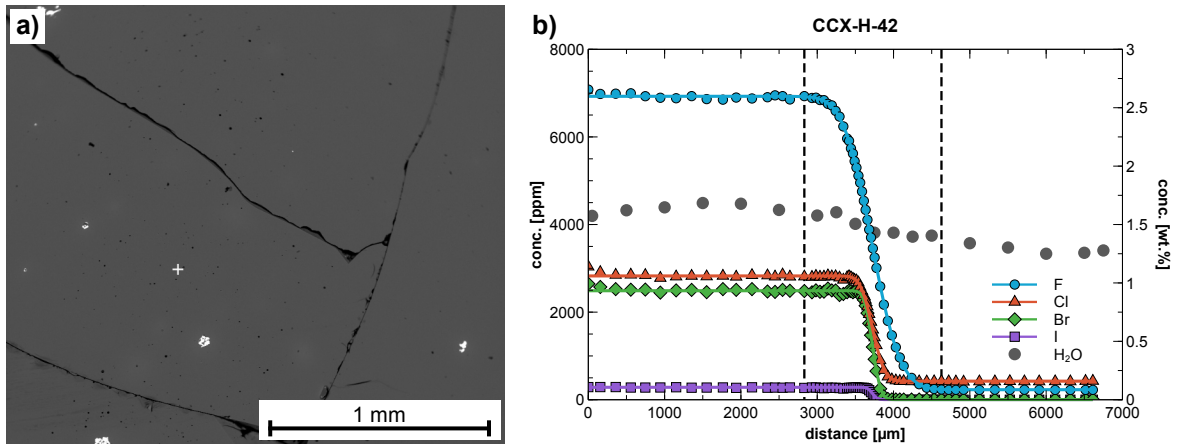


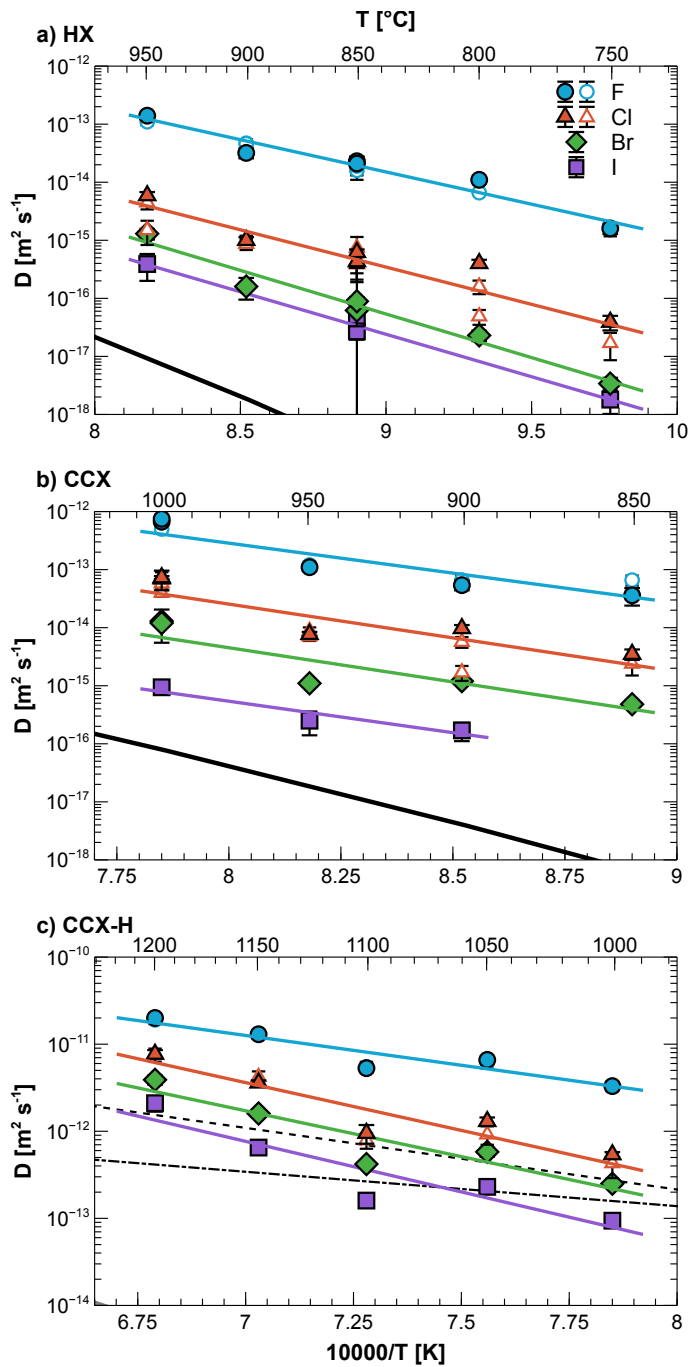
Fig. 1: Concentration vs. distance diagrams of one sample of each investigated melt composition acquired by EPMA and SIMS analyses. In the upper row, data of SiO₂ and totals refer to the secondary y-axis. EPMA data represents absolute values whereas the SIMS data was normalized to the average of the 4 outermost datapoints of each respective halogen in the enriched diffusion couple half. This way it is possible to visualize all halogens of a sample in one diagram.



936

937 **Fig. 2:** Backscattered image and diffusion profiles of sample CCX-H-42. **a)** Backscattered image of
 938 the center part of sample CCX-H-42 showing minor cracks likely caused during rapid quench (few
 939 seconds). **b)** Concentration vs. distance profiles of F, Cl, Br, I and H₂O as acquired by SIMS. The
 940 halogen data are plotted together with their respective fit curves calculated using equation 1.
 941 Diagrams of the other CCX-H samples are provided in the supplementary material.

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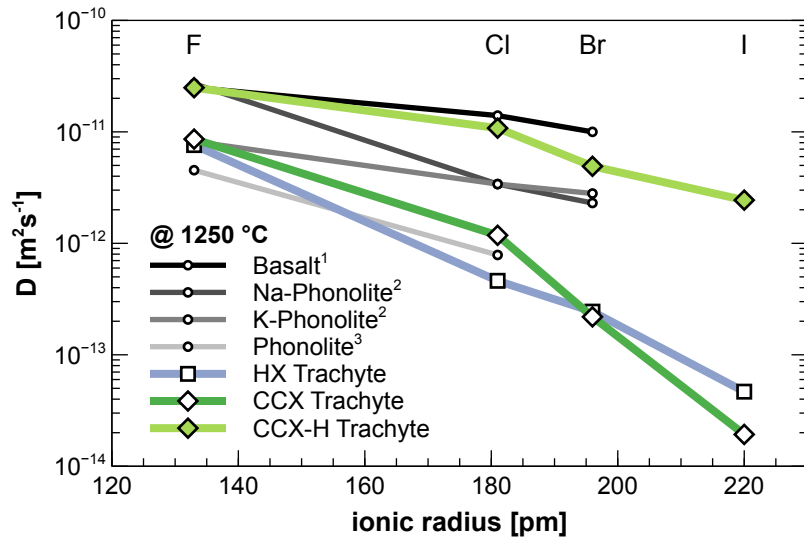


943

944 **Fig. 3:** Arrhenius diagrams of all investigated melt compositions illustrating the individual
 945 diffusion coefficients and the corresponding Arrhenius fits of all halogens. **a)** anhydrous HX melt.
 946 **b)** anhydrous CCX melt. **c)** hydrous CCX melt (~1.5 wt.% H₂O). The bold black lines indicate the self-
 947 diffusivity of oxygen calculated with the Eyring-equation and based on the calculated viscosities of
 948 the anhydrous melt compositions (Giordano et al., 2008). The dashed lines in c) indicate calculated
 949 oxygen self-diffusion in Rhyolite (dash-dot) and Dacite (dashed) from (Zhang and Ni, 2010). Open
 950 symbols represent data based on EPMA analyses.

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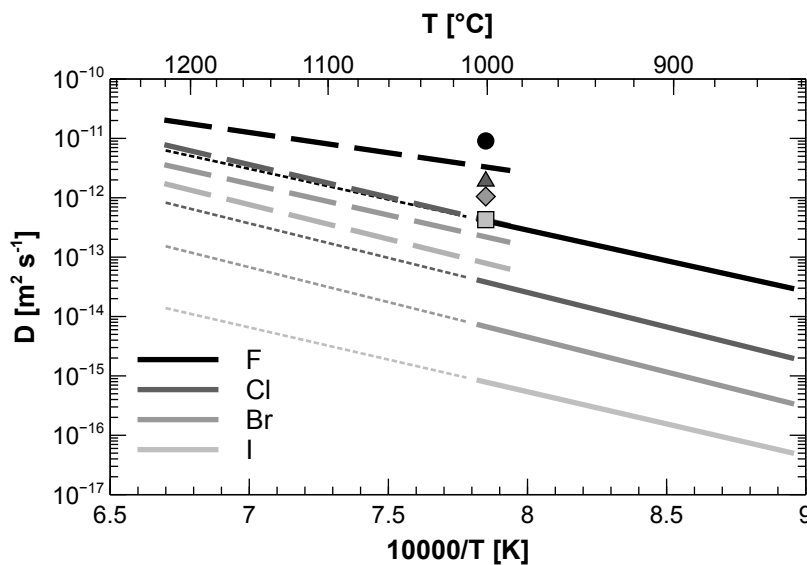
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953

954 **Fig. 4:** Halogen diffusivity at 1250 °C of the melts of this and other published studies plotted
 955 relative to the ionic radius of the respective halogen. All data are for anhydrous compositions
 956 except the hydrous data of the present study. Diffusivities at 1250 °C were calculated from the
 957 respective Arrhenius parameters of each composition and halogen. Data from: ¹ Alletti et al.,
 958 (2007); ² Balcone-Boissard et al., (2009) and Balcone-Boissard et al., (2020); ³ Böhm and
 959 Schmidt, (2013)

960



961

962 **Fig. 5:** Comparative Arrhenius diagram of halogen diffusion in anhydrous and hydrous (1.5 wt.%
 963 H₂O) trachytic melt determined in this study. Solid lines represent diffusion in anhydrous melt and
 964 dotted lines indicate extrapolation to higher temperatures. Long-dashed lines represent diffusion
 965 in hydrous melt. Solid symbols show preliminary results of halogen diffusion in melt containing 3.5
 966 wt.% H₂O. Note that the diffusivity of all halogens is even more enhanced in the more strongly
 967 hydrated melt.

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