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## 7                    **Fluid surface coverage showing the controls of rock** 8                    **mineralogy on the wetting state**

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### 16        **Key Points:**

- 17        • The analysis of fluid surface coverage is proposed as a novel approach to rock wet-  
18        tability characterisation
- 19        • A thermodynamically constrained model is derived and test on a Bentheimer sand-  
20        stone water-wet X-ray micro-CT dataset
- 21        • In a Berea sandstone, fluid surface coverage shows that rock mineralogy controls  
22        system local wettability after exposure to crude oil

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## Abstract

The wetting state is an important control on flow in subsurface multi fluid phase systems, e.g., carbon storage and oil production. Advances in X-ray imaging allow us to characterise the wetting state using imagery of fluid arrangement within the pores of rocks. We derived a model from equilibrium thermodynamics relating fluid coverage of rock surfaces to wettability and fluid saturation. The model reproduces the behaviour measured in a water-wet, nearly all-quartz, Bentheimer sandstone imaged during steady-state imbibition. A shift in fluid surface coverage is observed when the rock is altered to a new wetting state with crude oil. In two multi-mineralogical (Berea) samples, one water-wet and the other altered with crude oil, the analysis of fluid surface coverage after imbibition revealed mineral specific wetting preferences only in the altered system. Clays and calcite preferentially alter to an oil wet state, leading to mixed wettability in the rock.

## Plain Language Summary

When contacted by two or more fluids a solid surface may exhibit a preference for being coated by one of these fluids. This preference, called wetting preference, is crucial in defining how the fluids move in porous rocks or any other porous media of interest. The investigation of this wetting preference, defined wettability, is complex and several possible approaches are available in literature. In this study, we propose a novel approach based on the intuitive concept that the more a surface prefers to be coated by a certain fluid, the larger will be this coating, fixed the fluids' volume into the rock pores. The viability of our approach is first proven by considering two rock samples constituted by a unique material, but possessing different wetting preference. Eventually, we make use of our approach to better understand how in rocks constituted by a number of diverse materials, these materials behave in different ways when exposed to crude oil, mimicking the processes that happen in oil reservoirs. A deeper comprehension of this behaviour could aid the design of more efficient hydrocarbon production processes.

## 1 Introduction

Wettability is an important control in subsurface fluid flow, where fluids move through pore networks where capillary forces are dominant (Zou et al., 2018; Rücker et al., 2019; Lin et al., 2019). During oil recovery rock wettability exerts a control on the capillary entry pressure during primary drainage or in determining the likelihood of snap-off events of the non-wetting phase during waterflooding (Blunt et al., 2002). As a consequence of pore scale fluid dynamics, the behaviours of continuum scale properties such as relative permeability and capillary pressure are controlled by the wetting state (Anderson, 1987a, 1987b).

Predicting and characterising the wettability of an oil reservoir is a complex task. Minerals constituting rocks are naturally water-wet in the absence of hydrocarbon deposits. However, many oil reservoirs show relative permeability and capillary pressure functions indicative of intermediate-wet, mixed-wet or oil-wet systems (Donaldson et al., 1969). Indeed, rock surface wetting preference may be altered by the interaction of the solid substrate with surface-active compounds present in the crude oil. If present, these compounds can precipitate or diffuse to the solid surface and be adsorbed modifying the local wetting state (J. S. Buckley & Liu, 1998; J. S. Buckley, 1998). The results of these alteration mechanisms are dependent on the thermodynamic conditions, crude oil composition, brine composition and solid surface chemistry.

A number of studies have characterised the wetting behaviour of minerals typically found in the subsurface, i.e., in carbonate and sandstone reservoirs. A summary of the results of a collection of studies can be found in J. S. Buckley (1998). Calcite and clay minerals have been found to be more responsive to wettability alteration by crude oil

72 exposure than quartz (Alipour Tabrizy et al., 2011). However, experiments on chemi-  
 73 cally homogeneous flat surfaces or powders can only reproduce uniform altered wetta-  
 74 bility in the system considered. In order to investigate the role of rock geometrical com-  
 75 plexity and mineralogical heterogeneity in determining the *in situ* wetting state, it is nec-  
 76 cessary to study three-dimensional samples.

77 X-ray micro-CT offers the opportunity to investigate fluid arrangement inside rock  
 78 pores (Bultreys, Boone, et al., 2016; Bultreys, De Boever, & Cnudde, 2016; Coles et al.,  
 79 1996). With this newfound capability, thermodynamic theory indicates that it should  
 80 be possible to observe wetting signals from *in situ* contact angles, interfacial fluid cur-  
 81 vature and fluid-solid surface coverage (Morrow & Szabo, 1970). *In situ* contact angles  
 82 have been measured - either manually (Andrew et al., 2014; Singh et al., 2016) or au-  
 83 tomatically (Klise et al., 2016; Scanziani et al., 2017; AlRatrouf et al., 2017) - in the pore  
 84 space of various rock samples identifying different wetting states (Rücker et al., 2019;  
 85 Alhammedi et al., 2017). However, the measurements typically obtained have shown a  
 86 large variability in space and sensitivity to the processing pipeline chosen (Garfi et al.,  
 87 2019), making their direct employment difficult. Mean interfacial fluid curvature has suc-  
 88 cessfully been employed to map capillary pressure in water-wet and intermediate-wet rock  
 89 samples (Herring et al., 2017; Garing et al., 2017; Lin, Bijeljic, Pini, et al., 2018a; Lin  
 90 et al., 2019). However, the interpretation of mean interfacial curvature as a signal of wet-  
 91 ting is not straightforward: when the system is not water-wet, interfaces tend to have  
 92 null mean curvature, meaning that their curvature has opposite sign along the two prin-  
 93 cipal radii of curvature (Lin et al., 2019). Eventually, fluid-solid surface coverage as a  
 94 signal of wetting has not been explored thoroughly and its potential is still outstanding.

95 In this study, we show that the characterisation of fluid coverage of rock surfaces  
 96 can depict changes in the local wetting state. We develop and validate a simple model,  
 97 based in the thermodynamics of fluid-solid interfaces of a water wet system, to demon-  
 98 strate the applicability of solid surface coverage as a measure of wetting. Fluid-solid in-  
 99 terfacial areas are then measured to characterise the wetting state of two rock litholo-  
 100 gies. We first make use of observations on a mono-mineralogical rock (Bentheimer sand-  
 101 stone) as a case study to test the approach. We then extend our approach to chemically  
 102 heterogeneous systems and investigate mineral specific wettability in two Berea sand-  
 103 stone samples - one in its original state and one exposed to crude oil to alter the nat-  
 104 ural mineral wetting preference - by performing two drainage-waterflooding cycle exper-  
 105 iments and comparing the fluid arrangement observed in the two images acquired after  
 106 waterflooding.

## 107 2 Materials and Methods

### 108 2.1 Mono-mineralogical system: Bentheimer sandstone datasets

109 In this work we first make use of two datasets created by Lin, Bijeljic, Pini, et al.  
 110 (2018a) and Lin et al. (2019) as a case study with a simplified mineralogy. Bentheimer  
 111 sandstone is 98% quartz, 1% kaolinite/chlorite and 1% microcline, but for the purposes  
 112 of this work it was assumed to be a homogeneous rock constituted of a single mineral-  
 113 ogy. All the images were segmented into rock, brine and oil phases (Lin, Bijeljic, Pini,  
 114 et al., 2018a; Lin et al., 2019). In our study the region of interest used in the analysis  
 115 was  $900 \times 900 \times 3000$  voxels with  $3.58 \mu\text{m}$  voxel side, i.e. the spatial domain analysed  
 116 was  $3.22 \times 3.22 \times 10.74 \text{ mm}^3$ .

117 The first dataset - that we call Bentheimer Unaltered - consisted of the X-ray micro-  
 118 CT images acquired with two-fluid injection at five fractional flows ( $f_w = q_w/(q_w +$   
 119  $q_o)$  where  $q_i$  are volume flow rates of brine and oil) (S. Buckley & Leverett, 1942) of the  
 120 wetting phase, brine phase ( $f_w = \{0.15, 0.30, 0.50, 0.85, 1\}$ ), during steady-state imbi-  
 121 bition (brine fractional flow increasing with each step). The fluids in the system were

122 brine (3.5 wt% KI) and decalin. For further information please refer to Lin, Bijeljic, Pini,  
123 et al. (2018a).

124 The second dataset - that we call Bentheimer Altered - used a sample that was very  
125 similar to the Bentheimer Unaltered, except that the wetting state was altered before  
126 the coreflood. Prior to the flow experiments, this sample was partially saturated with  
127 crude oil and heated at 80°C for 30 days in a wetting alteration process known as age-  
128 ing. The fluids in this case were brine (3.5 wt% KI, 1.09 wt% NaCl, 0.02 wt% MgCl<sub>2</sub>.6H<sub>2</sub>O,  
129 0.11 wt% CaCl<sub>2</sub>.2H<sub>2</sub>O) and decalin (Lin et al., 2019; Lin, Bijeljic, Krevor, et al., 2018).  
130 Five images at fractional flow steps  $f_w = \{0.24, 0.50, 0.80, 0.90, 1\}$  were considered in  
131 this study.

## 132 **2.2 Multi-mineralogical system: experiments on Berea sandstone**

### 133 **2.2.1 Rock samples**

134 Two Berea sandstone samples of 4 mm in diameter and 20 mm in length were drilled  
135 from the same core. This core has laminations of cemented calcite. The main mineral  
136 groups present were identified by scanning electron microscopy (SEM) operated in back  
137 scattered electron (BSE) mode and coupled with energy-dispersive X-ray spectroscopy  
138 (EDS). Quartz grains constitute the majority of the rock matrix. The other mineral groups  
139 identified were clay group minerals (kaolinite, illite and smectite), potassium feldspar and  
140 small traces of minerals embedding metals. A reference example of mineral character-  
141 isation of Berea sandstone may be found in Lai et al. (2015). As with the Bentheimer,  
142 one of the samples was used unaltered by crude oil and is referred to as Berea Unaltered.  
143 The other sample underwent crude oil exposure after primary drainage and will be re-  
144 ferred to as Berea Altered.

### 145 **2.2.2 Fluids, fluid injection strategy, and wettability alteration**

146 Two drainage-imbibition cycle experiments were performed. In the experiment in-  
147 volving Berea Unaltered, the fluids employed were brine (15 wt% KI in de-ionized wa-  
148 ter) and decane. The sample was firstly saturated with brine at atmospheric pressure  
149 and then pressurized at the injection pressure of 3.5 MPa. Decane was thus injected at  
150 a flow rate of  $0.015 \frac{\text{ml}}{\text{min}}$ , which corresponds to a capillary number  $N_c \approx 10^{-7}$ . The to-  
151 tal injected volume of decane was 2.5 ml. The injection was stopped for at least 4 hours  
152 in addition to the scanning time, before performing brine injection. 40 pore volumes were  
153 injected at a constant flow rate of  $0.015 \frac{\text{ml}}{\text{min}}$ .

154 In the experiment with the sample Berea Altered, the fluids employed were brine  
155 (15 wt% KI, 1.09 wt% NaCl, 0.02 wt% MgCl<sub>2</sub>.6H<sub>2</sub>O, 0.11 wt% CaCl<sub>2</sub>.2H<sub>2</sub>O) and crude  
156 oil (density  $\rho = 0.8540 \text{ kg/m}^3$  and viscosity  $\mu = 4.7765 \text{ mPa s}$  at 20°C). The sample  
157 was firstly saturated with brine. Crude oil drainage was then performed by setting a con-  
158 stant pressure gradient of 5 Mpa between the injection and the receiving pumps, up to  
159 a total volume injection of 2.5 ml. After drainage, the sample was then removed from  
160 the coreholder and stored immersed in crude oil in a sealed glass bottle. The glass bot-  
161 tle was put into an oven at a temperature of 80°C for 30 days. After the wettability al-  
162 teration protocol, the sample was mounted in the coreholder and waterflooding was per-  
163 formed, by injecting 40 pore volumes of brine at a constant flow rate of  $0.015 \frac{\text{ml}}{\text{min}}$ .

164 In both the experiments, after waterflooding the injection was stopped and the sys-  
165 tem was allowed to equilibrate for 4 hours to a pressure of 3.5 MPa.

166 **2.2.3 Imaging and Image processing of Berea sandstone: minerals and**  
 167 **fluids phase segmentation**

168 The samples were imaged with an FEI Heliscan microCT obtaining a voxel reso-  
 169 lution of  $2.0\ \mu\text{m}$  for a region of interest larger than the sample cross section and a ver-  
 170 tical length of 8 mm. The projections were acquired while the sample was moving along  
 171 a helical trajectory and a 1 mm thick aluminium filter was employed. The X-ray source  
 172 voltage was set to 95 kV and the tube current to 70 mA. The raw images were reconstructed  
 173 employing an iterative back-projection algorithms provided by the scanner manufacturer.  
 174 For both samples, images were acquired before the injection of any fluid (referred to as  
 175 the dry scan) and after waterflooding.

176 The processing steps were the same for both samples. We filtered the dry scan and  
 177 waterflooding image by non-local means filtering (Buades et al., 2005) and registered them.  
 178 The greyscale dry scans were segmented using watershed segmentation (Beucher & Meyer,  
 179 1993) into five phases: pore space, clay group minerals, quartz-feldspar group minerals,  
 180 cemented calcite and others highly attenuating minerals. The filtered waterflooding im-  
 181 age was masked with the segmented pore space image, leading us to the segmentation  
 182 of the two fluid phases (oil phase and brine phase) by simple thresholding. The region  
 183 of interest of our analysis for each image was a cube of 1200 voxel side, i.e. 2.4 mm.

184 **2.3 Rock surface coverage as a measure of wetting: a model for water-**  
 185 **wet systems**

186 Consider a porous medium comprising two fluid phases, a wetting phase,  $w$ , and  
 187 a non-wetting phase,  $o$ , e.g., oil, and a solid phase,  $s$ . Per unit volume of pore space, the  
 188 reversible work required to increase the saturation of a non-wetting phase results in the  
 189 creation of fluid-fluid interfaces, between wetting phase, non-wetting phase, and the solid  
 190 (Morrow & Szabo, 1970; Bradford & Leij, 1997),

$$P_c dS_o = \sigma_{ow} dA_{ow} + \sigma_{os} dA_{os} + \sigma_{ws} dA_{ws} \quad (1)$$

191  $P_c$  is the capillary pressure,  $S_i$  is the saturation with  $S_o = 1 - S_w$ ,  $\sigma_{ij}$  is the interfa-  
 192 cial tension between fluid or solid phase  $i$  and phase  $j$ , and  $A$  is the interfacial area per  
 193 unit volume of pore space between phases.

194 The use of reversible work in the analysis is equivalent to limiting our considera-  
 195 tion to equilibrium states of the system, i.e.,  $P_c(S_o)$  and  $A_{ij}(S_o)$  at equilibrium. We ig-  
 196 nore irreversible work that may be required in practice to move from one state to the  
 197 next, e.g., due to transient processes (Berg et al., 2013; Morrow & Szabo, 1970). This  
 198 is the assumption made when making use of capillary pressure characteristic curves as  
 199 constitutive laws in the description of subsurface flow.

200 By integrating Eq.1, followed by algebraic operations and making use of the Laplace  
 201 relationship,  $P_c = 2\kappa\sigma_{ow}$ , where  $\kappa$  is the mean interfacial curvature of the oil-brine in-  
 202 terface, it is possible to derive the following (see the Supporting Information for a full  
 203 derivation):

$$A_{os}(S_o) = \frac{1}{\beta} \frac{\sigma_{ow}}{\sigma_{os} - \sigma_{ws}} \left( 2 \int_{S'=0}^{S_o} \kappa dS' - \int_{S'=0}^{S_o} \frac{dA_{ow}}{dS'} dS' \right) \quad (2)$$

204 The terms inside the brackets represent the reversible work of desaturation and the crea-  
 205 tion of oil-water interfacial area, respectively. The equation expresses the oil-solid in-  
 206 terfacial area created from the excess energy available when subtracting the work required  
 207 for the creation of fluid-fluid interfacial area from the work performed to increase the sat-  
 208 uration of the non-wetting phase in the rock. The multiplier term with the ratio of in-  
 209 terfacial tensions is equivalent to  $\frac{1}{\cos\theta}$  in a single capillary tube ( $\theta$  is the contact angle).  
 210 Without changing sign entirely, the more wetting the solid is with respect to the non-  
 211 wetting phase (the smaller the value of  $\sigma_{os}$ ), the more interfacial area between the non-

212 wetting phase and solid,  $A_{os}$ , will be created per unit of work. The parameter  $\beta$  repre-  
 213 sents a roughness factor that accounts for the mismatch between the real surface area  
 214 shared by each fluid and the solid surface and the one measurable by imaging, due to  
 215 imaging resolution limit (Helgeson et al., 1984; White & Peterson, 1990).

## 216 2.4 Rock surface coverage characterisation by micro-CT imaging

217 In order to characterise rock surface coverage, the interfaces between mineral phases  
 218 and fluid phases were identified. In the case of the Bentheimer datasets two groups of  
 219 interfaces were identified, between oil and rock, and between brine and rock phases. In  
 220 the case of the multi-mineral Berea sandstone, having produced segmented images with  
 221 four mineral phases and two fluids, we identified a total of eight interface groups, i.e.,  
 222 for each mineral and both fluid phases. Once an interface of interest was identified, a smooth  
 223 surface was constructed through that interface by means of a generalized marching cubes  
 224 algorithm.

225 We compare fluid surface coverage of different minerals by defining the fraction of  
 226 the total area of that mineral in contact with a fluid:

$$a_{i,j} = \frac{A_{i,j}}{\sum_i A_{i,j}} \quad (3)$$

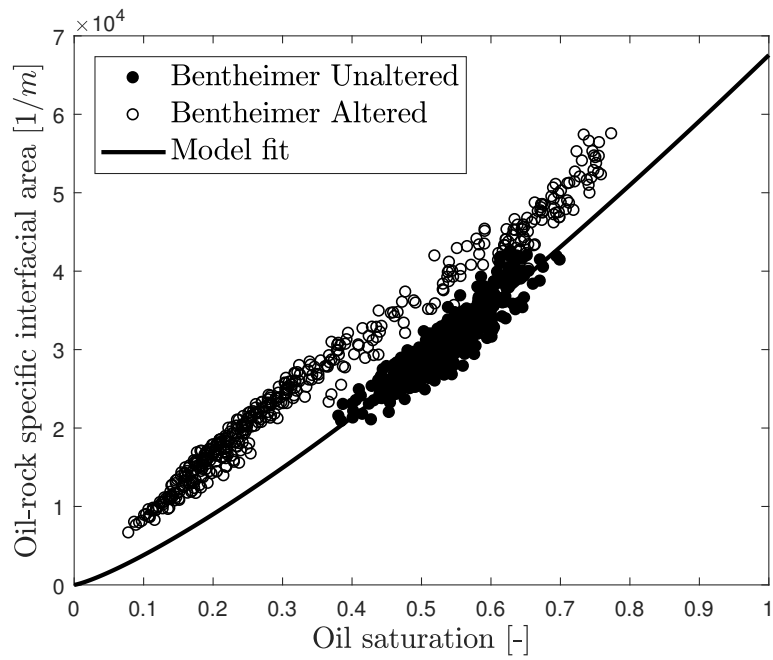
227 where  $A_{i,j}$  is the measured surface area per unit of pore volume shared by mineral  $i$  with  
 228 fluid  $j$ , respectively. The fractional definition of this property serves two purposes: to  
 229 allow for the comparison of the specific wetting preference of different mineral groups  
 230 with different total mineral-to-pore surface areas; to make the measurement more robust  
 231 to the surface smoothing and to the image processing pipeline chosen.

## 232 3 Results and Discussion

### 233 3.1 Bentheimer sandstone: fluid coverage of chemically homogeneous 234 rock surfaces

235 The region of interest for the ten images considered (five for Bentheimer Unaltered  
 236 and five for Bentheimer Altered) was divided into 90 cubic subvolumes of 300 voxels per  
 237 side (voxel side  $3.58 \mu\text{m}$ ). This allowed us to obtain a more precise topological descrip-  
 238 tion of the wetting state of the system investigated. In each of the subvolumes, for each  
 239 of the images and each of the datasets, fluid saturations, rock volume and fluid-coated  
 240 interfacial areas were computed.

241 The results obtained from the employment of our approach to wettability charac-  
 242 terisation reconciled well with the authors' assumptions that Bentheimer Unaltered was  
 243 water-wet (Lin, Bijeljic, Pini, et al., 2018a) and Bentheimer Altered was intermediate  
 244 or mixed wetting to oil (Lin et al., 2019). Specific oil-rock interfacial area measurements  
 245 are reported as a function of saturation for each of the subvolumes and for each of the  
 246 fractional flow considered in the two datasets Bentheimer Unaltered and Bentheimer Al-  
 247 tered in Figure 1. For similar oil saturation values, in the intermediate-wet - Bentheimer  
 248 Altered - sample rock coverage by oil is larger. As expected in a mixed or intermediate  
 249 wet system, oil is more likely to coat the solid surface than in a water-wet system. Fig-  
 250 ure 1 also shows that the model we proposed for water-wet systems (Eq.2) well repro-  
 251 duce the behaviour of the measured specific oil-rock interfacial area datapoints in Ben-  
 252 theimer Unaltered dataset. By fitting the model to the experimental data, we estimated  
 253  $\frac{1}{\beta} \frac{\sigma_{ow}}{\sigma_{os} - \sigma_{ws}} = 0.07$ . For a water-wet system  $\frac{\sigma_{ow}}{\sigma_{os} - \sigma_{ws}} \geq 1$ . This implies that the geo-  
 254 metrical roughness factor  $\beta \approx 10^1 - 10^3$ , consistent with literature roughness factor  
 255 values defined by comparing surface areas measured with BET to those estimated by X-  
 256 ray micro-CT imaging for other sandstone rocks (Lai et al., 2015). For additional infor-  
 257 mation on model fitting and the choice of input parameters, please refer to the Support-



**Figure 1.** Oil-rock specific surface area measured in two Bentheimer sandstone datasets consisting of the X-ray micro-CT images of two steady-state imbibition experiments. For similar oil saturation values, the oil-coated specific surface areas are larger in the altered sample than in the unaltered one. The behaviour of the experimental data defined for Bentheimer Unaltered is reproduced by the model (Eq.2)

**Table 1.** Mineral volumetric composition and remaining fluid saturation (oil remaining saturation  $S_{or}$ , brine remaining saturation  $S_{wr}$ ) after waterflooding from X-ray micro-CT images of the two Berea samples used in this study: Berea Unaltered (not aged by crude oil exposure) and Berea Altered (aged by crude oil exposure). The images were segmented in six phases: clay group minerals, quartz and feldspar, calcite cementation, other highly X-ray attenuating minerals, oil phase and brine phase.

	Berea Unaltered		Berea Altered	
	Mean [-]	St.Dev [-]	Mean [-]	St.Dev [-]
Clay	0.040	0.009	0.045	0.010
Quartz-Feldspar	0.840	0.046	0.813	0.043
Calcite	0.115	0.050	0.135	0.048
Others	0.006	0.004	0.006	0.004
$S_{or}$	0.572	0.050	0.239	0.071
$S_{wr}$	0.428	0.050	0.761	0.071

258 ing Information, where we also refer to Joekar-Niasar and Hassanizadeh (2012), Porter  
259 et al. (2009) and Raeesi et al. (2014).

260 Our methodology could be improved by the analysis of fluid coverage of rock sur-  
261 faces on a pore-by-pore basis, in order to increase the level of detail in the topological  
262 description of wettability.

### 263 **3.2 The role of rock surface mineralogy in controlling the wetting state**

264 In this case the region of interest was divided as the previous case in cubic subvol-  
265 umes of 300 voxels side, for a total of 64 subvolumes. In each of the subvolumes fluid sat-  
266 urations, mineral volume fractions and specific mineral fluid coating were computed.

#### 267 **3.2.1 Mineral composition and fluid saturation**

268 The segmentation of the images of the two Berea sandstone samples led to simi-  
269 lar mineral compositions (Table 1). This confirmed that the mineral segmentation work-  
270 flow employed is reproducible. The largest component of the rock matrix is the quartz-  
271 feldspar group minerals. Cemented calcite constitutes the second most abundant min-  
272 eral by volume fraction in the samples. Due to the process through which this cemen-  
273 tation likely formed, it is pore filling, exposing mineral surfaces only to poorly accessi-  
274 ble regions of the pore space. Segmented clay group minerals are broadly distributed,  
275 either as patches on quartz-feldspar or as clay aggregates.

276 The injection of 40 pore volumes of brine led to distinct values of remaining fluid  
277 saturation between the unaged and the aged samples. Berea Unaltered shows an aver-  
278 age remaining oil saturation of 57%, while in Berea Altered oil displacement was more  
279 effective, leading to an average oil saturation of only 24%. Mixed-wet conditions are more  
280 favourable to the recovery of the oil phase as has been observed extensively on larger core-  
281 flood tests (Salathiel, 1973). As observed for the Bentheimer datasets, the variability in  
282 saturation is larger for the sample that underwent the wettability alteration procedure.

#### 283 **3.2.2 Fluid arrangement in the pore space**

284 A visual inspection of the greyscale images acquired after waterflooding for the two  
285 samples clearly shows that the fluid arrangement differs. While in the unaltered sam-  
286 ple, clay minerals are mainly filled with brine after the waterflooding, in the aged sam-



287 ple, brine is prevented from invading the small pores of the clay (Figure 2). This is a qual-  
 288 itative signal that ageing has affected clay preferential wetting to brine.

289 As shown for Bentheimer sandstone, we expect oil-coated surface area fraction to  
 290 be positively correlated with oil saturation, i.e. the more oil in the pore space, the larger  
 291 the fraction of mineral surface area contacted by oil. However, for both Berea Unaltered  
 292 and Berea Altered this correlation is weak, as a consequence of the narrow range of fluid  
 293 saturation in the experiment.

294 The oil-coated surface area fractions computed for Berea Unaltered suggest that  
 295 all mineral groups considered are preferentially wetting to brine (Figure 3). Average oil-  
 296 coated surfaces are always less than water coated surfaces even at high oil saturation.  
 297 The average oil-coated clay surface area fraction is smaller than the quartz-feldspar frac-  
 298 tion and this may be due to pore or fluid morphology, or sub-resolution roughness. The  
 299 small pores found in these clays are preferentially imbibed by brine, due to the high cap-  
 300 illary pressure required for the non-wetting phase to occupy them. Similarly, calcite ce-  
 301 ment mainly exposes its surface area to brine. This is a consequence of the capillary pres-  
 302 sures associated with the narrow pore regions that the cementation did not clog when  
 303 it formed. These findings are consistent with previous studies identifying these miner-  
 304 als as water-wet. In the unaged sample pore geometry and rock texture are likely to be  
 305 responsible for the differences in the oil-coated mineral surface area fractions encoun-  
 306 tered. The system is uniformly water-wet.

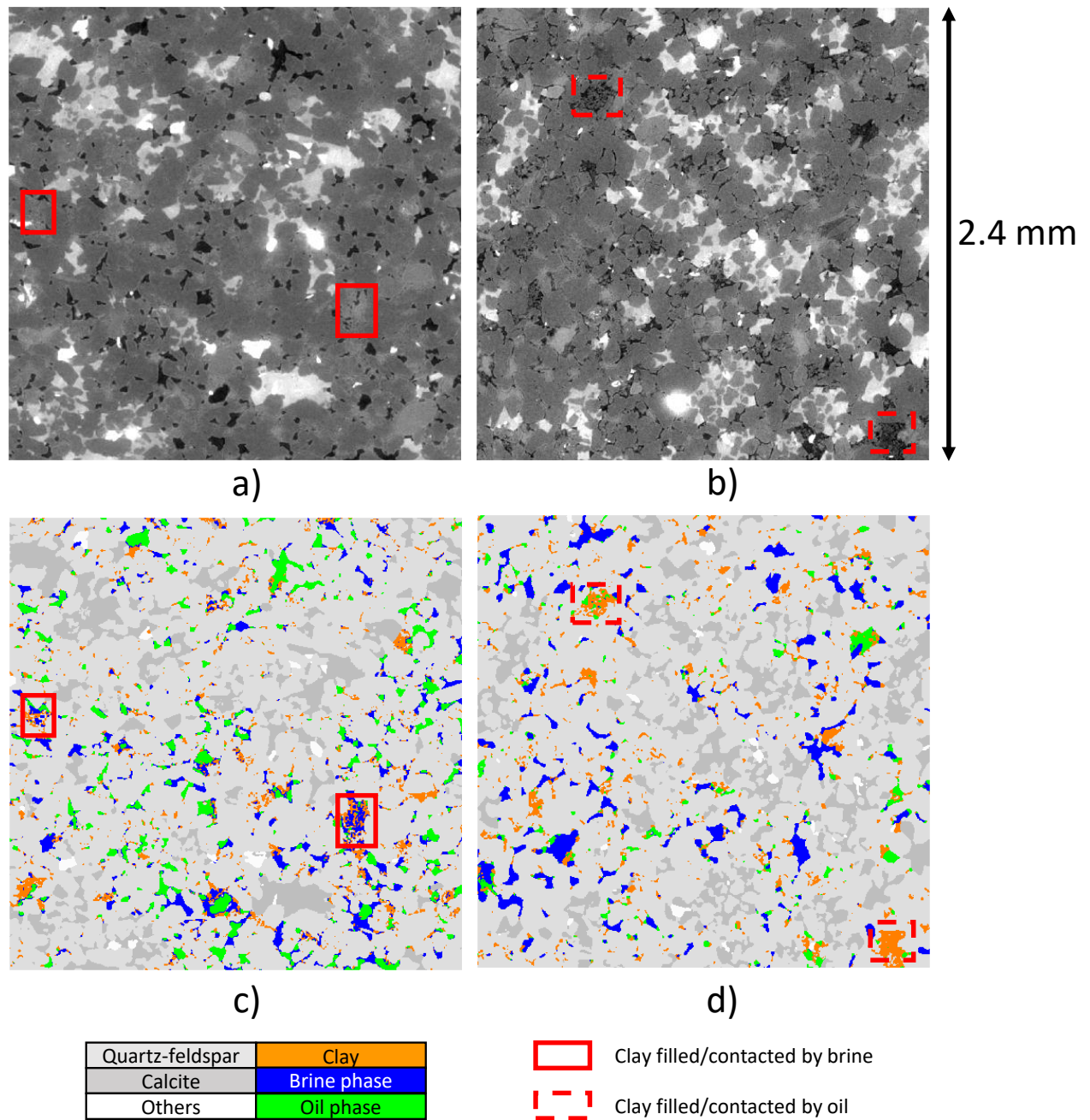
307 In contrast, rock mineral heterogeneities do control wettability alteration during  
 308 the ageing procedure. In Berea Altered, with a remaining oil saturation of 24%, 54% of  
 309 clay surface area is coated by oil. This shows a strong change in the wetting preference  
 310 of clay minerals, from water-wet to oil-wet. Even at lower oil saturation, there is much  
 311 higher surface area coverage of clay minerals by oil in the altered sample relative to the  
 312 unaltered sample. Similarly, a big increase is observed for cemented calcite, when results  
 313 for Berea Unaltered are compared to those obtained for Berea Altered. On the other hand,  
 314 quartz-feldspar does not show as strong of a wettability change. The reduced activity  
 315 of quartz-feldspar surfaces during ageing compared to those of clay and calcite is con-  
 316 sistent with what has been observed in Alipour Tabrizy et al. (2011), where clay and cal-  
 317 cite surfaces have been found more prone to wettability alteration.

318 The specific mineral behaviours we have identified in the altered sample suggest  
 319 that the wetting state is spatially correlated, with rock surface wetting preference chang-  
 320 ing with mineralogy. This may open up to the possibility of creating mixed wettability  
 321 maps based on mineral topological characterisation.

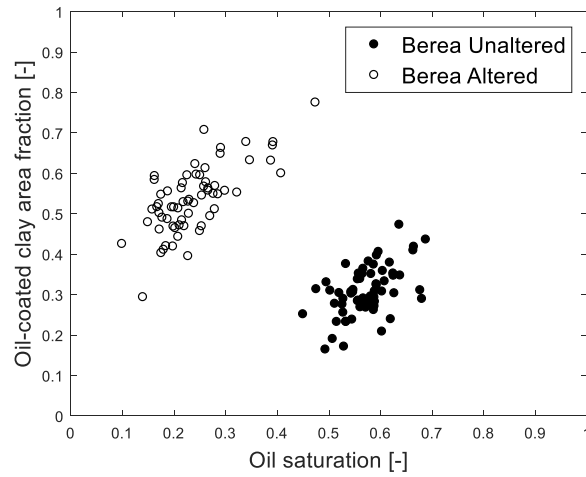
## 322 4 Conclusions

323 The analysis of rock mineral surface coverage by fluids can depict differences in the  
 324 wetting state of two fluid-phase systems. The solid surface covered by a fluid is positively  
 325 correlated with the saturation of that fluid. The particular relationship between fluid  
 326 saturation and fluid-mineral surface depends on the wetting state of the system. Con-  
 327 sidering the case of a uniformly water-wet system, we proposed a model that relates rock  
 328 coverage to fluid saturation, fluid-fluid interfacial curvature and fluid-fluid interface ex-  
 329 tent, measurements easily acquired with X-ray micro-CT imagery. This model was val-  
 330 idated by observations made before and after wetting alteration on a mineralogically ho-  
 331 mogeneous Bentheimer sandstone.

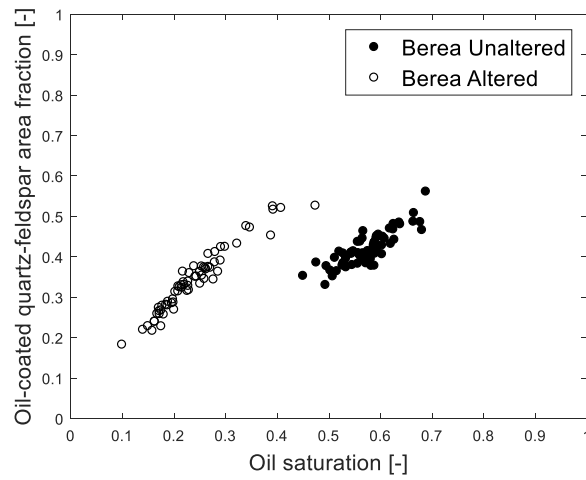
332 Rock surface coverage allowed us to investigate the role that mineralogy plays in  
 333 defining the wetting state of two sandstone rocks. In an untreated rock sample with sig-  
 334 nificant fractions of quartz, calcite, kaolinite and feldspar, fluid arrangement and surface  
 335 coverage after a drainage and imbibition displacement sequence were consistent with a  
 336 uniformly water-wet rock, regardless of local mineralogy. However, in a sample previ-



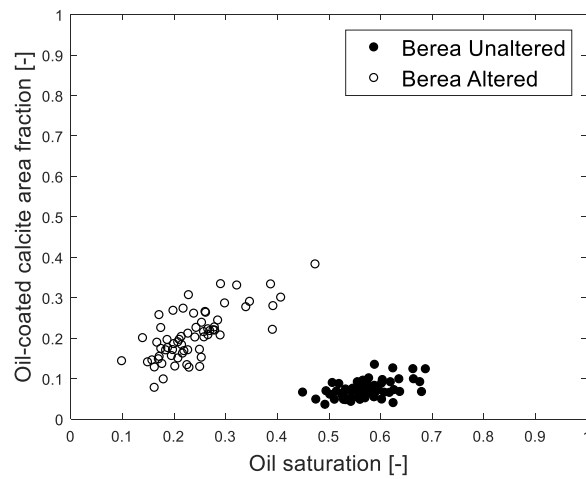
**Figure 2.** a) and b) show the greyscale images acquired after waterflooding of Berea Unaltered and Berea Altered, respectively. c) and d) show the segmented respective of a) and b). a) and b) show the change in clay wetting preference due to the effectiveness of the ageing protocol in Berea Unaltered and Berea Altered, respectively. In the sample Berea Unaltered clay aggregates are readily invaded by brine during waterflooding. In contrast, in the aged sample Berea Altered, brine invasion is largely prevented by the oil-wetting behaviour of clay surfaces.



a)



b)



c)

**Figure 3.** Oil-coated area fractions ( $a_{i,j}$ ) computed in the two Berea sandstone samples imaged after waterflooding. Quartz-feldspar group minerals show similar coating in the two samples. Instead, clay and calcite minerals preferentially altered to an oil-wet state, with an average increase in the oil-coated area fraction of 74% and 184%, respectively.

337 ously exposed to crude oil and high temperature for 30 days, mineralogical heterogene-  
 338 ity has been found responsible for heterogeneous wettability alteration processes. Clay  
 339 and calcite minerals were found more readily altered to an oil-wet state than quartz-feldspar  
 340 minerals. As a consequence, the sample wetting state was heterogeneous, mixed-wet, with  
 341 the distribution of the wetting state controlled by the local mineralogy.

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 345 jeljic, Pini, et al. (2018b). The Altered Bentheimer sandstone dataset and Berea sand-  
 346 stone datasets will be available at Digital Rocks Portal (<https://www.digitalrockportal.org/>),  
 347 currently underway).

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